

U.S. Department of the Interior  
U.S. Geological Survey

**Cruise Report,  
Hawaiian GLORIA Cruise  
F13-89-HW**

by

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## Summary of Scientific Results

1. The outer part of the southern Hawaiian EEZ between  $167^{\circ}\text{W}$  and  $179^{\circ}\text{E}$  longitude is characterized by relatively low and uniform acoustic reflectivity owing to the fairly subdued primary topographic relief in much of the area and also to the thick sediment cover that tends to blanket all but the steepest slopes.
2. Necker Ridge and three smaller ridges in the eastern part of the area are by far the most striking features in the survey area. Necker Ridge is oriented at about  $055^{\circ}$  and it appears to have been constructed by coalescence of many small circular cones. Several crossings of the ridge reveal a summit depression that is filled with 50-100 m of sediment; the summit depression extends along the ridge crest for at least 140 km.
3. The most common seamount in the area is a small (2-10 km diameter), circular and smooth-surfaced type that also is common in other Hawaiian EEZ regions. It typically has high reflectivity in a narrow band at the base of slope giving way to lower reflectivity in the center of the seamount. A summit crater can be seen in a few cases. Our 3.5 kHz profiles across several of these seamounts show that they are relatively smooth-surfaced, 300-500 m high, and their summit regions are in each case covered by 5-30 m of sediment. We suspect that the sediment cover is sufficient to absorb a part of the GLORIA signal and thus account for the lower reflectivity in the central area.
4. The other major seamount type in this EEZ is large (10-30 km across), rugged-appearing on GLORIA images, and non-circular. This type was especially common in a linear group of seamounts located at about  $23^{\circ}\text{N}$  between  $172^{\circ}\text{W}$  and  $173^{\circ}\text{W}$ . That seamount group appears to trend at about  $330^{\circ}$ , subparallel to the trend of the seafloor spreading fabric in this area and, intriguingly, also in alignment with the Line Islands trend, as shown for example by Karin Ridge, in the Johnston Island EEZ.
5. Three "lineations" consisting of ridges or seamount groups are shown on the GLORIA images about where the Murray Fracture Zone (MFZ) should cross the southern Hawaii EEZ. None of the lineations is particularly striking in this area owing to the thick sediment cover. Topographic

evidence for the MFZ is even less convincing closer to the Hawaiian Ridge, in the areas mapped during F1- and F2-90-HW.

6. The horst-graben fabric produced by seafloor spreading at the East Pacific Rise during the Cretaceous is clearly shown by the GLORIA data in the "central" part of the survey area (23-24°N, 175-176°W). The horsts are spaced about 2-5 km apart, they are sediment-covered, 50-100 m high, and their orientations range between 330° and 340°. This orientation is rotated 10-15° counterclockwise from the orientation observed in the F12-89-HW survey area north of the Hawaiian Ridge.

7. There is no evidence in this area for recent lava flows or large scale landslides like those found in the EEZ surrounding the eastern end of the Hawaiian Ridge.

8. Thick sections of sediment cover most of the area, attaining maximum thicknesses of 0.7 to 0.8 sec in basins near Necker Ridge. The average thickness in the area is on the order of 400 m.

9. An acoustically transparent layer up to 50 m thick is the most common bottom sediment type recorded by 3.5 kHz profiles. The transparent unit is amazingly clear and uniform in the western part of the EEZ, where it maintains a thickness of 20-25m along tracklines hundreds of kilometers in length. The sediment in this layer is thought to be pelagic and hemipelagic fine-grained material, and it commonly overlies discontinuously-stratified units that we infer to be turbidites, deposited when the adjacent Hawaiian Ridge was actively growing.

10. Although still extensive, the transparent "drape" becomes less continuous in the central part of the EEZ between 173° and 175°W longitudes. In that area the GLORIA imagery reveals an elaborate system of low relief channels that wind through the low ridges formed by the spreading fabric. Some of these channels are floored by acoustically-prolonged reflectors on the 3.5 kHz profiles suggesting that they have been recently-active conduits for turbidity currents.

11. Broadly-curving bands of low and moderate backscatter are present on the GLORIA images near a group of large seamounts at 23N, 173-174°W. The seismic data show that the bands are sedimentary features which we infer to be large mud(?) bedforms similar to those found by Dadisman and Marlow in the EEZ closer to the Hawaiian Ridge (F2-90-HW). However, in the present area there is very little topographic relief on these features and

we suspect that the reflectivity variations may be partly caused by subsurface variations in sediment properties.

## INTRODUCTION

Cruise F13-89-HW was the ninth GLORIA survey in a multi-year program designed to image the Hawaiian Island EEZ using GLORIA, a long-range side-looking sonar. The objective of this program is to produce atlases showing the geologic and morphologic features of the seafloor so as to better evaluate the economic potential, geologic hazards, and other possible uses of the Hawaiian EEZ.

Cruise F13-89-HW followed F12-89-HW which was carried out in October and November, 1989 (Torresan et al., 1991). F12 surveyed the area between Nihoa Island and St Rogatien Bank on the north side of the Hawaiian Ridge. This completed coverage for the first Hawaiian Gloria atlas which will extend from the eastern end of the Hawaiian EEZ to 167°W. In addition, F12 began coverage for the second Hawaiian Gloria atlas which will cover the remaining territorial waters out to the EEZ beyond Kure Island. Figures 1 and 2 provide geographic reference and trackline coverage for the area covered by this leg, F13-89-HW. The area surveyed is the outer part of the remaining territorial EEZ on the south side of the Hawaiian Ridge, tying in with the area surveyed by F10-88-HW (McGregor et al., 1989) on the eastern end at about 167°W and extending to the western end of the EEZ beyond Kure Island. The remaining area south of the ridge will be covered in two successive cruises starting in early January. These cruises will continue the pattern of long lines paralleling the ridge and extending the full length of the remaining EEZ survey area. The rest of the territorial waters of the Hawaiian EEZ will be imaged prior to 1992.

In addition to collecting GLORIA data, survey operations include two-channel seismic-reflection profiling using a 160 in<sup>3</sup> air-gun sound source, 3.5 kHz high-resolution profiling, 10 kHz bathymetric echosounding, magnetic and gravity field measurements, and upper water column temperature profiles using expendable bathythermographs (XBT).

## OPERATIONS

The GLORIA surveys are conducted from the M/V Farnella, a converted freezer-trawler that is under lease to the U.S. Geological Survey through the Institute for Oceanographic Sciences (IOS) in Wormley, England. GLORIA surveying responsibilities are split between USGS and IOS personnel. IOS personnel are responsible for all operations involving GLORIA, including developing the films and printing the images. IOS is also in charge of deck operations, maintenance of the seismic reflection and 3.5- and 10 kHz profiling systems and logging and final processing of navigation data. The USGS personnel are responsible for real time navigation, as well as for monitoring the gravimeter, magnetometer, seismic reflection, and the 3.5- and 10 kHz recording systems. The co-chief scientists (both USGS and IOS) are responsible for cruise planning, production of two field mosaics of the GLORIA data, and a preliminary science report. It is also part of the IOS chief scientist's duties to develop the films and print the GLORIA images.

The Farnella departed Honolulu, Hawaii on November 28, 1989. The scientific and ship personnel are listed below. The personnel list is followed by a schedule of field operations and a review of the equipment employed during the survey.

# **SCIENTIFIC STAFF FOR F13-89-HW**

## *U.S. Geological Survey*

LedaBeth G.Pickthorn	Co-Chief Scientist/Geologist
David E. Drake	Co-Chief Scientist/Geologist/
Michael Hamer	Navigator/Watchstander
Lisa Ramirez Bader	Geologist/DAFE/Watchstander
John Barber	Geologist/Watchstander
Michael Boyle	Electronic Technician

## *Institute of Oceanographic Sciences, U.K.*

Colin L. Jacobs	Co-Chief Scientist/
John Cherriman	GLORIA Chief Engineer
Robert Lloyd	Navigator
Alan Gray	Mechanical Engineer
Andrew Harris	Electronic Technician

## *JMarr (Ships Crew)*

John Cannan	Captain
Ronald Holliday	Chief Officer
Albert Fuller	2nd Officer
Michael Baldwin	Chief Engineer
David Rogerson	2nd Engineer
Roger Keys	3rd Engineer
Robin Searle	Electrician
Alan Thompson	Bosun
Jimmy Springall	Seaman
Peter Appleyard	Seaman
Michael Jessop	Seaman
Thomas Caughie	Chief Cook
David Graves	Steward

## Summary of Field Operations

The following list starts with the day of the year/Greenwich Mean Time (GMT), for the starting point of the major survey segments of F13-89-HW . When converting to local time note that the day 332 is November 28, and that GMT is 10 hours ahead of local Hawaii time, eg., 1900 GMT is 0900 local. Figure 1 provides a geographic reference for the region covered by this survey and figure 2 is a trackline summary to provide a reference for the various stages of the survey.

### Day 332 (Tuesday, November 28)

- 0000 - Took a gravity reading at the gravity base station at the Bishop Museum in Honolulu with a portable gravity meter for the purpose of establishing a gravity tie..
- 0820 - Made a land-sea gravity tie at pier 13 with the portable gravity meter.
- 1910 - Sailed from Pier 13, Honolulu.
- 2212 - Began logging navigation and gravity on ABC system.

### Day 334 (Wednesday, November 29)

- 0735 - Arrived at launch point. Calm seas, windy and drizzling.
- 0800 - 3.5 and 10 khz bathymetry fish deployed..
- 0900 - Gloria and magnetometer deployed.
- 1000 - Hydrophone streamer deployed.
- 1005 - Began line 1, course 287°, all equipment working fine. Continue through line 5, 338/1650, at which point we decided to do the western ends of our two longest lines while the weather was favorable since it would have been more likely to encounter bad weather on that end.

## Day 338 (Monday December 4)

- 1705 - We turned north along line 6, then got back on to course  $288^{\circ}$  on line 7. We surveyed out to the western end and came back along line 9 to the same approximate latitude at which we started line 7. This finished off the western end of the survey area.

## Day 339(Tuesday, December 5)

- 2200 - Turned southward along line 8 and picked up the southern part of the survey where we left off.

## Day 349(Saturday, December 16)

- 2004 - Ended line 17, did the short tie line, 18 and started line 19, tying in to lines 7 and 9, done earlier in the cruise.

## Day 352(Monday, December 18)

- 1805 - Ended line 19, completing assigned survey. At 1820 began a tie line (line 20) to connect with the first line assigned to the January GLORIA survey.
- 1956 - Started line 21, beginning the new survey area.

## Day 353(Tuesday, December 19)

- 0639; - Ended line 21 prematurely. Due to 35 knot winds the GLORIA engineer decided to recover GLORIA, not wanting to take the chance of hazardous recovery conditions the next day when we were scheduled to make the recovery. He was concerned that if the high winds were to persist through the night it would produce a heavy swell, making recovery difficult and dangerous.
- 0800 - The airgun, streamer, magnetometer and GLORIA, were recovered, and all but GLORIA were then redeployed for an airgun survey to make use of our leftover time.



- 0807    -    Began airgun survey (line 22) on a course of 034° to make an airgun tie line roughly orthogonal to the long lines in the next two GLORIA surveys, with the plan of going toward the Hawaiian ridge to a depth of 2000 m. We planned to cross what looked like a landslide on the Chase bathymetry.
- 0918    -    Owing to time becoming short, we changed course to 051, toward another landslide target that didn't look as promising, but was closer. We were not sure of being able to complete the line all the way across the landslide before it was time to haul in the gear for the transit back to Honolulu.
- 2159    -    Finished line 23, a very short line designed to spend our extra time surveying along the ridge until it was time to begin our transit back to Honolulu.
- 2230    -    Seismic gear, magnetometer, and both bathymetry fish were recovered and we began our transit to Honolulu.

### Equipment Summary

This section summarizes problems encountered with the shipboard data collection systems. Appendix I (from Torresan et al., 1989) summarizes the standard operational procedures that were established for the 1986 surveys (Holmes et al., 1987; Normark et al., 1987 and 1989). Complete reviews of the trouble-shooting and repairs for each system are available in the electrical technicians' report.

#### *Gravity Meter*

The gravity meter, a LaCoste and Romberg S-53, functioned continuously for the entire cruise. A land-sea gravity tie was

established at pier 13 in Honolulu prior to departure. This was done by taking a reading on the dock and at the IGSN reference gravity base station at the Bishop Museum in Honolulu with a LaCoste and Romberg model G portable gravity meter number 426. Another land-sea gravity tie was established following the cruise so that a drift correction could be applied to the data. Began logging gravity on magnetic tape immediately after leaving port and began logging it on the ABC system at 2200 the same day.

### *Magnetometer*

The magnetometer was deployed following the deployment of GLORIA (approximately 334/0840). Official logging began at 334/0900. The data are recorded on both strip-chart and magnetic tape, with the gravity data being recorded on the same magnetic tape. There were no problems with the system. At 351/2200 about 40 minutes of noise was noticed on the strip chart. The magnetometer was then pulled in and it was found to have some cloth (the ET said "someone's knickers") wrapped around it. This was removed and the magnetometer was redeployed by 2230. The data collected during the noisy period were still usable so the data gap was only the 30 minutes which it took to test that the fish was operating properly. The magnetometer was recovered prior to recovering the GLORIA fish on about 353/0645 and redeployed by 0800 for use during the airgun survey. Final recovery was at 353/2100.

### *Expendable Bathythermographs (XBT's)*

The XBT probes were deployed once daily, beginning on JD 332, to measure the thickness and temperature of the surface mixed layer, and the temperature profile within the thermocline. T-4 probes are capable of profiling to 460 m, and T-7 XBT's can profile to 760 m. T-7 probes were deployed twice on the trip and T-4 probes were used the rest of the time. The system consists of an XBT launcher and receiver, a GOES satellite transmitter, and a micro-

computer that handles recording, plotting, formatting and data transmission. The system performed well, with some fouling of the wire in the towed magnetometer. The probe was launched using a 4 foot extension handle on the port side of the vessel, about 20 feet forward of the stern. The following list is a record of the location of daily XBT drops and their success rate.

Day	XBT	Depth	Rating	Record Length	Latitude	Longitude
332	T-04	460 m	Full		21 12.3'N	158 40.2'W
333	T-04	460 m	Full		20 32.3'N	163 52.3'W
334	T-04	460 m	Partial	(298m)	20 50.9'N	168 00.1'W
335	T-04	460 m	Partial	(307m)	21 11.4'N	167 20.0'W
336	T-04	460 m	Full		22 12.2'N	170 42.4'W
337	T-04	460 m	Full		23 16.2'N	174 18.9'W
338	T-04	460 m	Full		24 39.9'N	177 09.8'W
340	T-04	460 m	Full		26 00.0'N	178 54.7'E
340	T-04	460 m	Partial	(343m)	25 13.1'N	178 05.9'W
341	T-04	460 m	Full		23 10.2'N	175 50.8'W
342	T-04	460 m	Full		22 24.9'N	172 22.5'W
343	T-04	460 m	Full		23 36.0'N	176 15.5'W
344	T-04	460 m	Full		23 47.9'N	175 10.7'W
345	T-04	460 m	Full		22 40.4'N	171 21.4'W
346	T-04	460 m	Full		21 27.6'N	167 18.7'W
347	T-04	460 m	Full		22 27.8'N	169 42.0'W
348	T-07	760 m	Full		23 34.5'N	173 30.1'W
349	T-04	460 m	Full		24 32.0'N	176 48.0'W
350	T-04	460 m	Partial	(379m)	23 54.0'N	173 43.2'W
351	T-07	760 m	Partial	(365m)	22 46.6'N	169 51.8'W
352	T-04	460 m	Full		21 54.1'N	166 35.0'W
353	T-04	460 m	Full		23 37.2'N	166 45.6'W

### *3.5-kHz High-Resolution Profiling System*

The 3.5-kHz reflection tow fish was deployed on 334/0800. The system was operational and official logging commenced immediately. The system performed well with only routine maintenance throughout the course of the cruise. Generally, record quality degrades in water depths greater than 4800 m when steaming at speeds in excess of 8.5 kts. We generally averaged about 8.5 kts and therefore were able to collect good quality records to depths exceeding 5000 m. The tow fish was recovered 353/2230.

### *10-kHz Echo-Sounding System*

The 10-kHz echo-sounding system employs a tow fish similar to that of the 3.5-kHz system, and was deployed on 334/0800. The system worked well, with down time restricted to routine maintenance, and blade and paper roll changes. The tow fish was recovered 353/2230.

### *Two-Channel Seismic Reflection System*

The two-channel seismic reflection system employs a 2600 cm<sup>3</sup> (160 in<sup>3</sup>) air-gun sound source, and an 800 m two-channel streamer (including a weighted stabilizing section). Two 50 m long active sections are towed about 500 to 600 m behind the air gun. The air gun is fired every 10 seconds. Two channel data is recorded on a MASSCOMP computer, and a 2.3-second analog record of one of the channels is displayed on a graphics monitor. In addition, two single channel analog hard copy records are produced on Raytheon line scan recorders (LSR), one 8 seconds and one 6 seconds. The memory function of the LSR printing the six second record was used to print the profile with a constant orientation (west and north ends of profiles are on the left, east and south ends are on the right) Vertical exaggeration is about 4:1 at a speed of about 8.5 knots on the 6 second record and 21:1 on the 8 second record.

The two-channel seismic reflection system performed well during the survey. The streamer and air gun were deployed by 334/1000, and official logging began immediately. The only system down time, besides routine airgun swaps performed for the purpose of maintainance such as replacing seals, was one 41 minute period when the airgun went down and was swapped (348/2033 to 348/2114). Once for about 2 hours (349/1526 to 349/1733) the MASSCOMP hung up, resulting in a 2 hour gap in the digital data on the tape. A system reboot corrected the situation. This problem occurs roughly twice per trip. Programming changes need to be made to avoid this in the future. The airgun and streamer were recovered prior to recovery of the GLORIA fish on 353/0700 and redeployed for an airgun survey by 0800. Final recovery was at 353/2200.

### *Shipboard Positioning Systems*

The navigation system (described in Holmes et al., 1986 and Normark et al., 1987) consists of a navigation program that runs a real time graphic steering display using an IBM PC XT. Two monitors show this display, one is on the bridge where it is used for steering and one is in the lab where progress is monitored and navigation inputs are selected. This program processes Loran signals from a Northstar 7000 receiver to do direct ranging on Loran stations (Rhorho navigation), and also receives Global Positioning System (GPS) fixes from a Trimble 4000 receiver. GPS is the most accurate and fixes were available for about 16 hours per day in a single long interval. Rhorho was selected when GPS was not available, but it was unusable much of the time, due to periods of drift which appeared to be artificial. For example, after a period of tracking a relatively straight course along the line, the navigation display would suddenly show the ship's position drifting, usually northward, with course over ground differing from the ship's heading by up to 11 or 12 degrees, when normally it is only a few degrees. We adopted a policy of telling the bridge to maintain the course when these drifts would occur, ignoring the display. We would then use transit satellite fixes, received on the bridge, to decide what course to steer until GPS was

available. Sometimes the drift was subtle and we relied on the bridge to make this judgement. They did an excellent job in this respect due to their experience and attentiveness. This policy worked fairly well, in that we generally maintained a reasonable course and often ended up within several hundred meters of the GPS position when it became available. On our final navigation plots upon which the GLORIA mosaic was made, we used dead reckoning based on the ship's speed and steering, for times when no acceptable navigation was available. One curious thing is that rho-rho tracked quite well in the far western and eastern portions of the survey area, with little or no drift problem. We were using the Central Pacific Loran chain with the master on Johnston Island, and slaves on Hawaii and Kure Islands. The geometry was really no better in these areas than in the central portion of the mapping area.

### *GLORIA Side-scan Sonar System*

The deployment and operation of the GLORIA system is covered in extensive logs by the IOS personnel. More detailed summaries are available in Somers et al. (1978), Laughton (1981), and references therein. A summary of the GLORIA pass record and the number of files is presented in Appendix 2. Note that one pass equals 6 hours.

The GLORIA system was launched on about 334/0900 and operated continuously without problems until 353/0630 when the survey was terminated. To avoid spurious echo returns, signal transmission was suspended during course changes. The GLORIA recovery, as usual followed that of the magnetometer and seismic gear.

### *GLORIA Shipboard Image Processing*

The techniques employed in shipboard processing are described in detail in Normark et al. (1987 and 1989) and will not be elaborated upon here. No special shading techniques, such as those used by Normark et al.(1989), were applied. Following the printing

of the GLORIA images, the images were laid down over corrected and smoothed navigation plots and mosaicked. Clear overlays were made, upon which geological features were mapped out and the thicknesses of soft sediment were recorded.

## RESULTS

For the purposes of consistency and ease of comparison with earlier reports we will present this report as a series of topical sections guided by the format and results obtained during previous GLORIA surveys of the Hawaiian EEZ. The following outline organizes the topics addressed:

1. Cretaceous seafloor spreading fabric
2. Murray Fracture Zone
3. Seamount distribution and characteristics
4. Cretaceous ridges
5. Sediments
  - a. 3.5 kHz echo character
  - b. Sediment distribution and GLORIA backscatter intensities
  - c. Channel systems
  - d. Bedform fields
  - e. Summary of sedimentary features and GLORIA backscatter.

### Cretaceous Seafloor Spreading Fabric

Seafloor spreading fabric is subdued, owing to the thick accumulations of pelagic mud, often referred to as drape, that is pervasive throughout the F13-89-HW survey area. On the GLORIA image this subdued fabric manifests itself as faint lineations between 5 and 40 km. in length (fig. 3). The fabric has an orientation that ranges between 330° and 340° and it is present in two areas. One area is at 25°N, between 178° and 179°W. The other is between 23° and 24°N, and between 175° and 176°W. In both of these areas,

where the lineations cross the ship's track, the 3.5 kHz bathymetry record shows small sediment-covered ridges that are from about 10 to 100 m in height. Some of these ridges are fairly symmetrical but most exhibit an asymmetrical cross section, with significantly more relief on one side than the other (fig. 4). The thick sediment layer that has nearly buried the Cretaceous spreading fabric topography is well illustrated by the air gun record in figure 4.

The GLORIA data show another area of interesting "lineations" located at about 24°N and 174°W. Their orientations are generally similar to the spreading fabric direction, but the lineations themselves are slightly curved. Examination of 3.5 kHz profiles across these features shows that they are sedimentary channels, with orientations which are influenced by the structural grain of the horst and graben spreading fabric. These channels can be traced across multiple tracklines and are described in more detail in a subsequent section of this report.

The spreading fabric lineations on the GLORIA sonographs are often lines across which very subtle changes in acoustic backscatter take place. These changes in acoustic backscatter may be produced by increases in sediment compaction over buried ridges or due to topographically-produced backscatter variation (Kayen et al., 1990).

### Murray Fracture Zone

There are three nearly linear low ridges or groups of seamounts in the area that trend roughly perpendicular to the trend of the spreading fabric. We consider these to be possible extensions of the Murray Fracture Zone. One of these features is pair of low linear ridges separated by a 25 km wide band of spreading fabric topography occurs between 24° and 25°N and between 178° and 177°W. The orientation of the band is about 065° and the length is about 110 km. Interestingly, this lineation does not appear to line up with any similar features in the area to the north, mapped by cruise F1-90-HW (Kayen et al., 1990). However, this pair of low ridges and the intervening subdued horst-graben elements appear to be roughly



located on the trend of the Murray Fracture Zone as mapped by Torresan, Clague and Jacobs during F12-89-HW on the northern side of the Hawaiian Ridge. (Torresan et al.,1991).

The second lineation in our survey area consists of a line of very small seamounts (a few km in diameter), possibly extending to a larger seamount (6 km. across) at the northeast end. This line of small hills and seamounts crosses latitude  $24^{\circ}\text{N}$ , between  $175^{\circ} 30'$  and  $176^{\circ} 30'\text{W}$ . Its orientation is between  $070^{\circ}$  and  $075^{\circ}$ . The feature either dies out before reaching the EEZ region closer to the Hawaiian Ridge or is too deeply buried to be detected by GLORIA.

The third lineation consists of a group of en echelon(?) short ridges or fractures. located at  $24^{\circ} 40'\text{N}$  and  $175^{\circ} 30'\text{W}$ , near the northern edge of our survey. These elements are slightly offset to the southeast from lineations that are interpreted to be part of the Murray Fracture Zone in the F1-90-HW area to the north (Kayen et al, 1990). The "fractures" at the northern edge of our area have high acoustic backscatter on the GLORIA image but are short and disjointed, making it difficult to determine an exact orientation. Inspection of the 3.5 kHz record on line 19 (350/0700) shows a rugged seamount 400-450 m in height that is covered by 5-10 m of sediment.

### Seamount Distribution and Characteristics

Most of the seamounts imaged on this cruise are less than 10 km in diameter. One common morphology revealed by the GLORIA images is circular with high acoustic backscatter around the base of the seamount and lower backscatter in the central area (fig 5). 3.5 kHz profiles across two of these seamounts, which are 6 and 7.5 km in diameter, show them to be about 375 and 450 m high, respectively, smooth surfaced with steep basal slopes and gentler upper slopes which are covered with 10 to 30 m. of sediment (fig. 6). The sediment covering probably accounts for the low GLORIA backscatter of the central portions of these seamounts, and differing basal and upper flank slopes may account to a lesser degree for the

contrasting acoustic backscatter on the GLORIA images. The 3.5 kHz profiles also show possible "moats" at the bases of the two seamounts, 10 to 20 m. deep which appear to exist because of thinner sediment deposits than in the adjacent areas farther from the seamounts. The airgun profile shows the height of these two seamounts over acoustic basement is between 700 and 800 m. There are roughly 30 of these circular seamounts scattered throughout the area mostly in groups; for example, the group at the far western end of the survey area. In addition, some of these circular seamounts occur in the midst of the smaller ridges to the west of Necker Ridge. The geometry of the seamounts in that region suggests that the ridges may be composite features, built by a number of coalesced, single-vent, circular seamounts (fig. 7).

A few of the small seamounts appear non-circular and have a rough texture on the GLORIA image, giving them the appearance of being composite features. We didn't cross directly over any of these, thus we lack 3.5 kHz profiles to help interpret the GLORIA images.

Most of the remaining small seamounts are so small that it is hard to recognize any details of their morphology. They occur as bright spots on the GLORIA image and are small enough that if the ship had passed directly over one, it would be obscured on the GLORIA image by the nadir, thus it is difficult to match GLORIA images of these with 3.5 kHz crossings.

As discussed earlier, two possible lineations made up of the circular and the very small seamounts occur in the general vicinity of the Murray Fracture Zone as mapped by Kayen et al.(1990), and are roughly parallel to it. These may have been, as discussed by Fornari, et al.(1984) and Torresan et al.(1991), formed at the mid-ocean ridge spreading center, contemporaneously with the formation of the Cretaceous ocean crust.

The larger seamounts in our survey area range from 20 to 36 km. across and tend to be non-circular. There are seven in the area, five of which occur in a rough line, trending about 330°, between 171° and 173°W. A crossing over the summit of one seamount about 27 km wide shows relief on the order of 3200 m. These larger seamounts appear rugged and mottled on the GLORIA images, with

areas of high and moderate backscatter. In places they appear to be composite features, possibly made up of smaller cones and vents (fig. 7). However, the 3.5 kHz summit crossing mentioned above does not support that observation. It shows the seamount slope to be very steep but also coherent. Nevertheless, the echo character of the slope does vary on the 8-sec. air-gun record (fig. 8), possibly accounting for the variations in the GLORIA backscatter. 3.5 kHz profiles of the lower slopes of the large seamounts show them to be rugged, varying from being suggestive of deep gullying and intense degradation to almost complete obliteration of the data by hyperbolic returns (fig. 9).

### Cretaceous Ridges

The most striking feature in the area is Necker Ridge, which cuts across the area at an orientation of about  $055^\circ$  between  $167^\circ$  and  $169^\circ\text{W}$  (see fig. 2). The ridge is over 500 km long and extends from Necker Island on the Hawaiian Ridge southwestward to the Midpacific Mountains. The image of Necker Ridge on our GLORIA sonograph mosaic appears as a chain of flat-topped seamounts. These "flat tops" are artifacts that result from gaps in the "far-field" GLORIA coverage. GLORIA swath width is strongly dependent on water depth, becoming narrower as the depth decreases. Since the majority of our mapping area is deeper than 4500 m, the basic trackline spacings were calculated accordingly. Consequently, on oblique crossings of shallow features like Necker Ridge the insonification coverage is narrower, leaving the outer edges of the sonograph dark. Mosaicking these passes together results in a chain of "holes" in the coverage such as we have described. These data gaps are unacceptable, and they were filled on two later cruises, F13-90-CP and F1-91-CP, which each insonified one side of Necker Ridge on their transits out to the Johnston Island EEZ. Each paralleled the crest, offset enough to provide coverage in the upslope data voids.

During F13-89-HW, we made seven crossings of Necker Ridge which showed the relief to be between 2000 m in the south and

3000 m in the north (fig. 10). Necker Ridge appears very rugged on the GLORIA sonographs, with generally high acoustic backscatter. Patchy dark areas of low reflectivity along the summit area and on the ridge flanks may represent shadowing of the acoustic beam and perhaps some ponds of low-reflectivity sediment. Features which cannot be extrapolated across the trackline are difficult to interpret because the airgun and 3.5 kHz systems have relatively narrow beam widths corresponding with the areas obliterated by the nadirs on the GLORIA sonographs. Sediment ponds on or near the summit of Necker Ridge were observed on the airgun records of three crossings, lines 15, 17 and 21 (fig. 10). Two of these sediment ponds matched up well with a band of uniformly lower acoustic reflectivity, insonified by the later cruises, which appears to follow the summit area between latitudes  $21^{\circ} 25'$  and  $22^{\circ}\text{N}$ . A similar, smaller area occurs on the GLORIA sonographs between latitudes  $22^{\circ}10'$  and  $22^{\circ} 30'$ , but no 3.5 kHz profiles of it exist because no crossings were made there.

The GLORIA mosaic constructed during F13-89-HW shows areas of intermediate acoustic backscatter on each side of the base of the ridge. By "intermediate" we mean lower backscatter than the bedrock slopes of the ridge itself but higher than the surrounding basin areas, which are covered by thick pelagic sediment. We interpret these ridge basal areas to be accumulations of relatively coarse-grained debris that has been shed from the ridge slopes. Airgun records and 3.5 kHz profiles show rugged foothills that may represent alluvial fans as well as larger slope failures. We note, however, that there is little evidence of the sort of large landslide deposits that are so common close to the relatively young, actively-growing and degrading, eastern end of the Hawaiian Ridge (Clague and Dalrymple, 1987; Lipman et al., 1988; Moore et al., 1990). The GLORIA image of Necker Ridge acquired on our line 15 clearly shows a number of gullies and small canyons trending directly down the sides of the ridge attesting to the degradation of this edifice. But large slumps, debris slide deposits or landslide "scars" are lacking.

Four other smaller elongated ridges occur in the vicinity of Necker Ridge, two on the east side to  $166^{\circ}\text{W}$ , and two on the west to

169° 30' W. These and several linearly arranged groups of seamounts are subparallel to Necker Ridge, with orientations of 055° to 065°. The smaller ridges appear on the GLORIA image to be composite features, especially to the north where the small circular variety of seamounts, described earlier, seem to be superimposed on one another within the structure of the ridges. The ridge which is farthest to the west (at about 169°W) goes from 2500 m. of relief at our most southerly crossing on line 5 to 1300 m on line 17 to the north. Further north it disintegrates into a wide area of these small circular seamounts, that in some cases have central craters (fig. 7).

In a few places along Necker Ridge the GLORIA image shows features that appear to be individual cones or overlapping flows. We hypothesize that all of these ridges may have formed through the buildup of relatively small individual cones and volcanoes, with the most highly-developed ones becoming so massive that the individual elements are now obscured. It is unclear whether this buildup would be progressive or occur more or less simultaneously along the ridge length. The trend of these ridges is roughly parallel to the Murray Fracture Zone and roughly perpendicular to the Cretaceous spreading fabric. Thus, it seems possible that Necker Ridge might have formed as a "leaky transform fault", which could probably have either progressive or simultaneous vent eruptions. The variation in trend of these ridges (10-15 degrees) casts doubt on this as a mode of formation for all of them. Radiometric ages on Necker Ridge, published in the literature to date, are very few, and in the authors' opinions, unreliable. Detailed sampling of these ridges needs to be done, to carry out a thorough geochronological study. This would definitely shed light on the formational history of these unique volcanic edifices.

## Sediments

### *a. 3.5 kHz echo character and sediment distribution*

We recognize three major groups of 3.5 kHz bottom reflectivity type in this region: (1) hyperbolic reflectors, (2) smooth opaque

reflectors, (3) acoustically transparent reflectors. These categories and our descriptions of the echo types differ somewhat from those set up and discussed by Normark et al. (1987) and Torresan et al. (1991). This may be a result of differences in sediment types and their distributions in our respective areas. It is important to point out that the groups are somewhat simplistic subdivisions of a gradational spectrum of echo types. We decided it was best to keep our groupings general because some small variations in the echo character can actually be due to variable equipment settings, ship's speed or interference by other equipment.

We refer to hyperbolic reflections in this paper to mean overlapping hyperbolic-shaped echoes with no subbottom reflections. Deeper reflectors are therefore either absent or masked by the hyperbolic echoes. Normark et al. (1987) and Torresan et al. (1991) recognized this echo type and found it to be typical of rough topography consisting of either rugged bedrock ridges or blocky slumps and landslide deposits. The size of the hyperbolae is probably directly correlatable to the size of the roughness, ranging from surface roughness on a lava flow to giant blocks from a landslide. Hyperbolic echoes can occur buried beneath transparent drape as well as when cropping out on the ocean bottom (Fig 11).

The second type, opaque reflectors, consist of non-hyperbolic echoes which mask any subbottom stratification or structure. They are sharp at the top boundary and vary from a thin band to a thick fuzzy band of disorganized reflections (referred to as "prolonged"). This echo type is thought to be produced by hard sediment or bedrock surfaces. The "prolonged" nature of the echo may be due to small scale roughness on surfaces that produce a multitude of pinpoint returns of variable length (Wallin, 1982). This results in the fuzzy shading we see on the 3.5 records that fade out with depth. These opaque reflectors are one of the most common types in the area, occurring most often as subbottom reflectors beneath the transparent layer, but also have been observed on the slopes of many of the Cretaceous seamounts and ridges in the area as well as in channel floors (Figs 12 and 14).

The last 3.5 kHz bottom type consists of a reflector which is acoustically transparent (Fig. 12). This can cap acoustically clear layers that are thought to be made up of uniformly fine-grained pelagic or hemipelagic mud (Wallin, 1982; Normark et al., 1987). A clear transparent blanket ranging from a few meters up to nearly 50 meters in thickness is one of the most common acoustic units in this region. It is occasionally observed ponded on some seamount peaks and slopes, but in general it is present throughout the region on gentle slopes and in flat areas, grading from clear to stratified in many areas. A clear transparent sediment layer is especially widespread on the gentle slope west of the MFZ. The layer forms a "pelagic drape" for hundreds of kilometers along our two western survey lines, 7 and 9 (Fig. 2). We believe it is especially prevalent in that region because sources of terrigenous sediment are remote. If channel systems ever existed in the western area they must have been abandoned and buried long ago (Torresan et al., 1990). The uniformity of the thickness of the transparent drape in the western area also gives evidence that the bottom currents are weak there.

The transparent bottom reflector can also cap units which contain internal stratification (acoustic laminations), consisting of one or more discrete reflectors separating transparent sediment units (Fig. 12). Individual reflectors are laterally discontinuous, extending from less than a kilometer, to tens of kilometers. This stratified transparent layer can also grade laterally into a clear transparent layer. The discrete reflectors are inferred to be composed of coarser sediments that are episodically deposited, such as volcanic ash or turbidite deposits. The reflectors vary from sharp to prolonged in character, sometimes laterally grading from one to the other. The 3.5 kHz records in the F13-89-HW area show that the stratified-transparent sediment layer occurs near large seamounts and near the ridges in the eastern part of the survey area. In addition, we observe stratified-transparent units associated with an extensive channel system that crosses the very gently-sloping seafloor between the longitudes of 173°W and 175°W (Fig 13). This channel system is described later in this report.

*b. Sediment distribution and GLORIA backscatter intensities*

GLORIA backscatter intensity is a complex function of sea floor slope, and the composition and surface roughness of the ocean floor. Ground truthing experiments carried out by IOS and the USGS Branch of Marine Geology in the past, have been devoted to deciphering the complex relationships between these factors and the corresponding GLORIA backscatter intensities (J. Gardner, personal communication, 1990). Recently, it has become clear that the GLORIA signal can penetrate into sediment covered bottoms to variable depths of up to 5-20m, further complicating the interpretation of backscatter targets (J. Gardner and M. Hampton, personal communication, 1990).

They have found that because of these complexities it is difficult to make generalizations that would help in deciphering these relationships based only on the GLORIA images during a survey. Whereas volcanic bedrock on seamounts and ridges produces obvious intense acoustic returns, sediment type variations produce much more subtle backscatter variations that can be similar to the acoustic signals from small-scale topography and bottom roughness changes (eg. sedimentary bedforms). However, there are some rather clear correlations between 3.5 kHz echo type and the GLORIA backscatter levels from sediment-covered areas in this survey region.

Survey lines 4,5, and 15 (see figure 2) provide several good examples of these correlations. Beginning in the eastern area around Necker Ridge, there is a good correlation between moderately high backscatter areas along the bases of these ridges and stratified-transparent and opaque sediment bottom types on our 3.5 kHz records. Careful inspection of the sonographs in this eastern area also reveals a correlation between thick transparent sediment layers and low GLORIA backscatter. This is in contrast to the relatively strong backscatter from flat channel floors that exhibit a smooth opaque 3.5 kHz reflection (fig. 13).

Subtle backscatter variations occur over the gentle sediment-covered slopes in the western part of the survey area. Nevertheless, it is apparent on lines 5 and 15 that sea floor regions that are uniformly blanketed by a clear-transparent layer of 15-20 m or more are also uniformly darker on GLORIA images. Conversely, in the



same areas, decreases in sediment thickness and/or the presence of stratified-transparent layers appear to cause a subtle increase in GLORIA backscatter(as would be expected).

A final note of caution regarding these correlations is in order. The depth of penetration of the Gloria sound pulse in an area covered by an acoustically-transparent sediment is still uncertain and it is likely that some backscatter patterns are caused by subsurface features.

### *c. Channel systems*

The GLORIA sonographs and 3.5 kHz data reveal individual channels and channel systems in several parts of the survey area. This is in contrast to the results of previous Hawaiian EEZ surveys along the younger eastern part of the island chain (Normark et al., 1987) in which channel systems were not common. We find evidence for both inactive and recently active density-current channels in the area. The channels are quite wide, ranging from less than a km in width to as much as 5 km (fig. 13). However, none are incised more than a few tens of meters below the surrounding sea floor. The channel floor deposits vary from acoustically-opaque to clear-transparent overlying either stratified-transparent or opaque reflectors(fig. 12c,d). Our preliminary interpretation is that valleys and channels that exhibit an opaque echo character on relatively gentle slopes are active, or were recently active because no detectable pelagic drape has accumulated in the channel. Channels with opaque reflectors occur parallel to the base of the slope on the east side of Necker Ridge and on the west side of Nihoa Ridge. While we cannot determine the full extent of these channels until the next two survey legs are completed we are inclined to believe that they have been formed by localized slumping of sediments down the steep flanks of the adjacent ridges followed by the generation of turbidity currents. As mentioned above, the base-of-slope channels near the Cretaceous ridges return strong opaque acoustic signals which we assume to be caused by coarse volcanogenic sediment.

A more extensive channel system that we believe will ultimately be shown to have begun on the Hawaiian Ridge is

revealed by the GLORIA sonographs between 173° and 175°W (Fig. 13b,c). The seafloor in this region slopes very gently to the south and the topographic relief is generally less than 50 meters. The GLORIA data show an intricate system of dark and light broadly-curving and meandering streaks and bands that cross this surface from north to south. These represent channels, the orientations of which appear to be controlled by the Cretaceous horst and graben spreading fabric. We believe that this sediment transport system has been partially abandoned and is presently inactive because the channels have been partially or completely filled with acoustically-transparent sediment.

*d. Bedform fields*

Two areas on either side of a prominent chain of seamounts located at 23°N, 173° 30'W show arcuate bands of alternating high and low backscatter on the sonographs (Fig 13a). The smoothly curving nature of these bands and their presence on a depositional slope of very low relief suggests that they are large mud(?) bedforms. The wavelength of these features ranges from about 3 km near the seamounts to less than 1 km as one proceeds downslope to the southwest. The field covers an area of about 400 km and appears to converge toward the seamount chain. This geometry suggests that the bedforms could be the result of currents flowing through a gap(?) between the seamounts. However, the origin of these backscatter features is certainly not clear at this point. Although it would be interesting to attempt to "ground-truth" these features, our 3.5kHz data, shown in figure 14b, suggests the possibility that the GLORIA reflectivity was controlled by bedforms that are now rather deeply buried by pelagic drape.

*e. Summary of sedimentary features and GLORIA backscatter.*

It is apparent that many of the backscatter patterns in this area are related to sediment transport processes, facies changes as shown by 3.5 kHz echo type, and deposition that was significantly influenced by currents (Fig. 12d). Some of the backscatter variation is due to local changes in sediment type at the bases of seamounts and ridges, where the present-day bottom currents may be

accelerated by the topography. Other channel systems are areally extensive and are likely to be the offshore parts of turbidity-current transport systems that originated on the Hawaiian Ridge. An understanding of the history of these features would shed additional light on the growth and degradation of the island chain. Finally, there is no question that the GLORIA reflectivity in this part of the Hawaiian EEZ is generally subdued relative to the images acquired to the east near the younger part of the ridge and the latest images collected in the Johnston Island and Palmyra Island EEZ's. The subdued reflectivity is at least partly due to the thick and extensive "transparent" layer found in this survey.

### ACKNOWLEDGEMENTS

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## FIGURE CAPTIONS

**Figure 1.** Location map of the Hawaiian EEZ showing the area surveyed by cruise F13-89-HW.

**Figure 2.** Trackline chart for F13-89-HW (Mercator projection).

**Figure 3.** GLORIA image of the spreading fabric present on line 11.

**Figure 4.** Airgun profile (8-sec full scale) over the spreading fabric shown in figure 3.

**Figure 5.** GLORIA images of small seamounts. a) circular, smooth-surface types. b) small non-circular seamounts.

**Figure 6.** 3.5 kHz profile over a small, circular seamount showing the steep lower slopes and the sediment-covered summit region.

**Figure 7.** GLORIA images of large, irregular seamounts.

**Figure 8.** Airgun profile over a large seamount. Note the changes in reflection strength on the seamount slopes which we believe may produce the "rough-surfaced" quality in the GLORIA data shown in figure 7.

**Figure 9.** 3.5 kHz and airgun profiles over the lower slope of a large, gullied, seamount (see figure 7).

**Figure 10.** Airgun profile across Necker Ridge showing the sediment-filled, summit depression and the possible slump(?) deposits at the base of slope.

**Figure 11.** 3.5 kHz profile showing a clear transparent layer over a rough subsurface reflector with hyperbolics.

**Figure 12.** Examples of the main 3.5 kHz echo types in this area. a) acoustically-transparent over opaque-prolonged. b) stratified-transparent. c) transparent over discontinuously-stratified. d) an example of the irregular distribution of the transparent layer within a region of low-relief turbidity-current(?) channels (see figure 13).

**Figure 13.** GLORIA images of (a) an inferred bedform field and (b,c) low-relief channel systems in the central part of the survey area.

**Figure 14.** An example of prolonged acoustic reflectivity on the 3.5 kHz profile (upper panel), and a portion of the 3.5 kHz record over the inferred bedform field shown in figure 13a (lower panel).

**APPENDIX I****EQUIPMENT SETTINGS AND COMMENTS*****3.5 kHz SYSTEM***

LSR Recorder	Mode - Continuous PAPER - 100 lpi SWEEP - 1 sec PROGRAM - As required GAIN - Mid CONTRAST - Mid THRESHOLD - Min
--------------	---

PTR Transceiver	GAIN - 6 POWER - -6db PULSE WIDTH - Not used
-----------------	--

IOS Correlator	OUTPUT LEVEL - 4-5 ATTENUATOR - 11.5
----------------	---

Fish Depth Compensation 10 m.

***10 KHz SYSTEM***

MUFAX Recorder	ATTENUATOR - -6 to -18 TIME MARKS - 6 min. PULSE LENGTH - 2.8 to 5 FISH DEPTH - 2x5 m GATING SELECTOR - 6 FATHOMS/METERS - Meters TRIGGER (Left/Center) -as required for scale changes TVG - Use to suppress outgoing pulse GATING - Use to "see" bottom through outgoing pulse.
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***MAGNETOMETER***

Soltec Recorder	CHART SPEED - 40 cm/hr All three channels at 100 gamma range
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***GRAVIMETER***

Soltec Recorder	CHART SPEED - 10 cm/hr
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***SEISMIC REFLECTION SYSTEM***

LSR Recorder	DISPLAY - Normal STYLUS SCAN - 2 sec PAPER - 120 lpi
--------------	--

MEMORY SWEEP - 6 sec  
FILTER - out  
POLARITY - +  
GAIN 10 to 2 O'clock  
CONTRAST -30 to -40  
THRESHOLD - 2

Krohn - Hite Filter  
Masscomp

20-120 Hz  
GAIN SETTING: Pre-amp 0  
Post-amp 6 to 18  
As required for DWD

DMD Delay Box



**APPENDIX 2****GLORIA PASS RECORD**

<b>Pass</b>	<b>Start Time</b>	<b>End Time</b>
1	334/0900	334/1459
2	334/1500	334/1559
3	334/1559	334/2059
4	334/2100	335/0259
5	335/0300	335/0459
6	335/0500	335/0659
7	335/0659	335/1015
8	335/1016	335/1559
9	335/1600	335/2159
10	335/2200	336/0359
11	336/0400	336/0959
12	336/1000	336/1559
13	336/1600	336/2159
14	336/2200	337/0359
15	337/0400	337/0959
16	337/1000	337/1559
17	337/1600	337/2159
18	337/2200	338/0359
19	338/0400	338/0959
20	338/1000	338/1559
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26	339/2200	340/0359
27	340/0400	340/0959
28	340/1000	340/1559
29	340/1600	340/2159
30	340/2200	341/0359
31	341/0400	341/0959
32	341/1000	341/1559
33	341/1600	341/2159
34	341/2200	342/0359
35	342/0400	342/0959
36	342/1000	342/1559
37	342/1600	342/2159
38	342/2200	343/0359

Pass	Start Time	End Time
39	343/0400	343/0829
40	343/0829	343/1359
41	343/1400	343/1959
42	343/2000	344/0159
43	344/0200	344/0759
44	344/0800	344/1359
45	344/1400	344/1959
46	344/2000	345/0159
47	345/0200	345/0759
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78	352/2000	353/0159
79	353/0200	353/0639

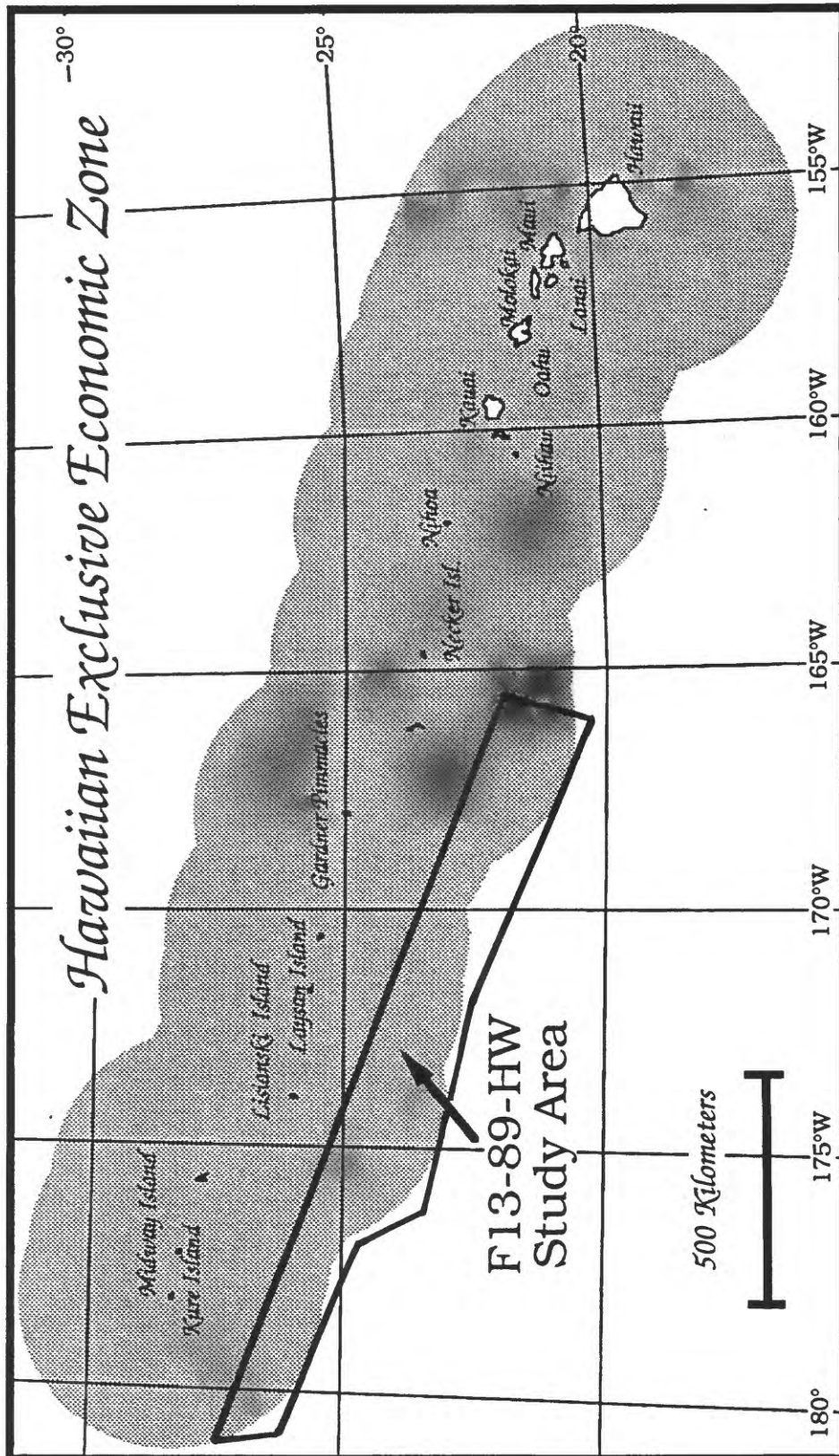


Figure 1. Location map of the Hawaiian Island EEZ and study area of F13-89-HW.

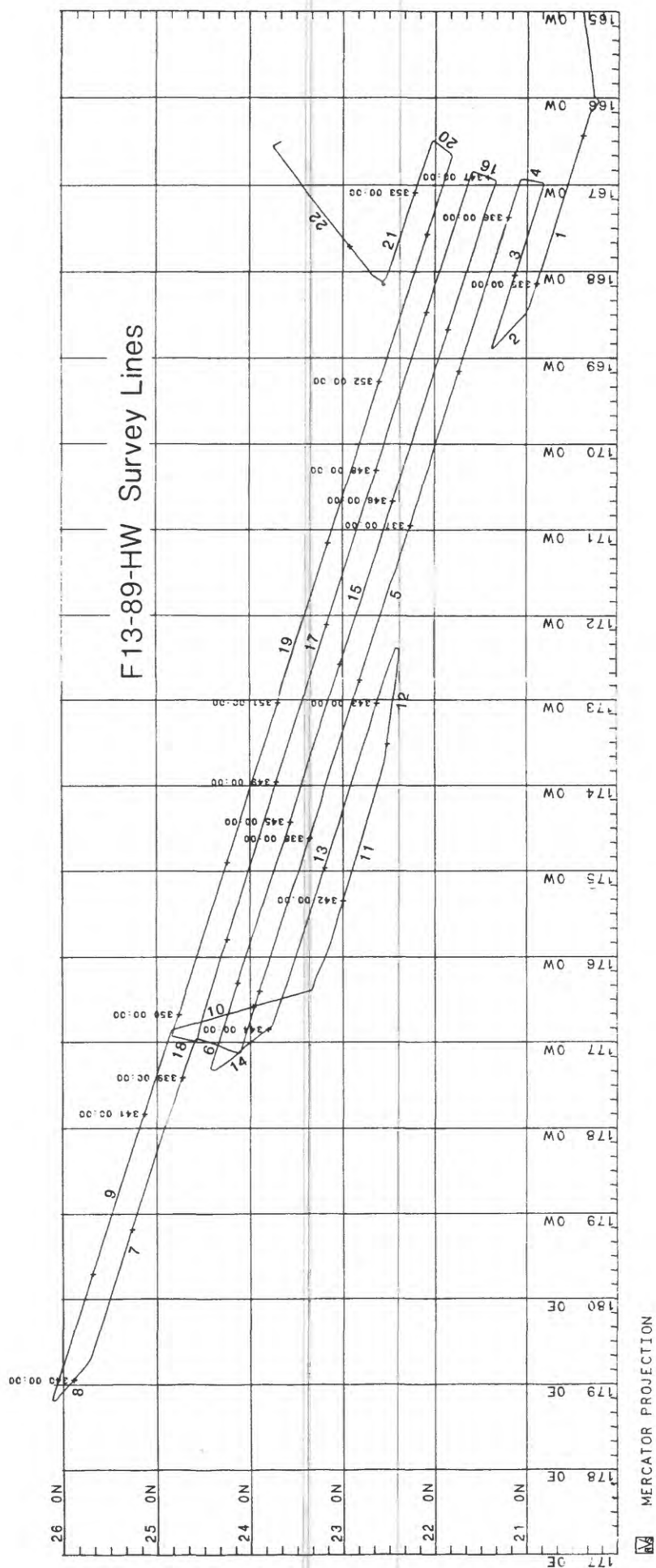


Figure 2

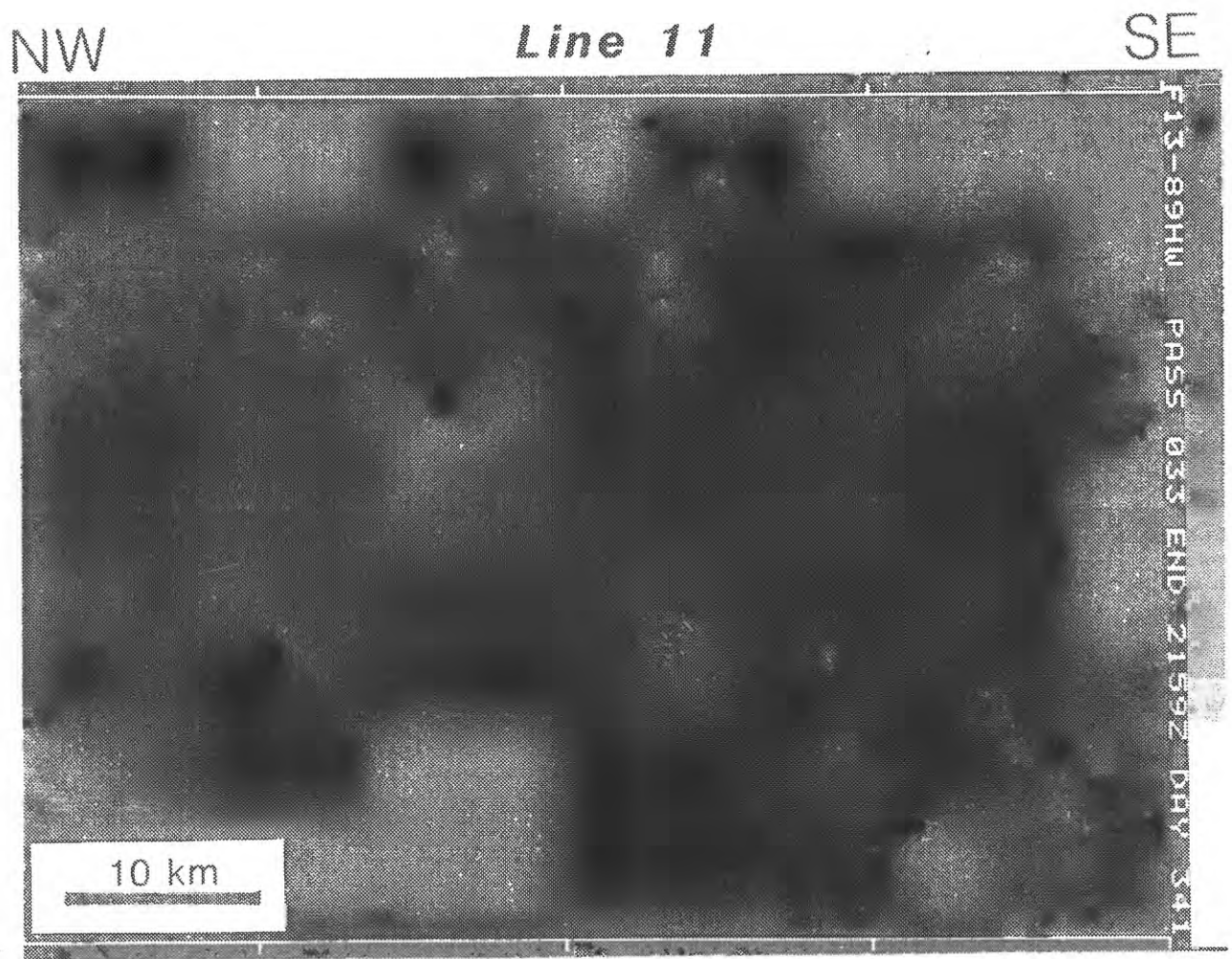


Figure 3

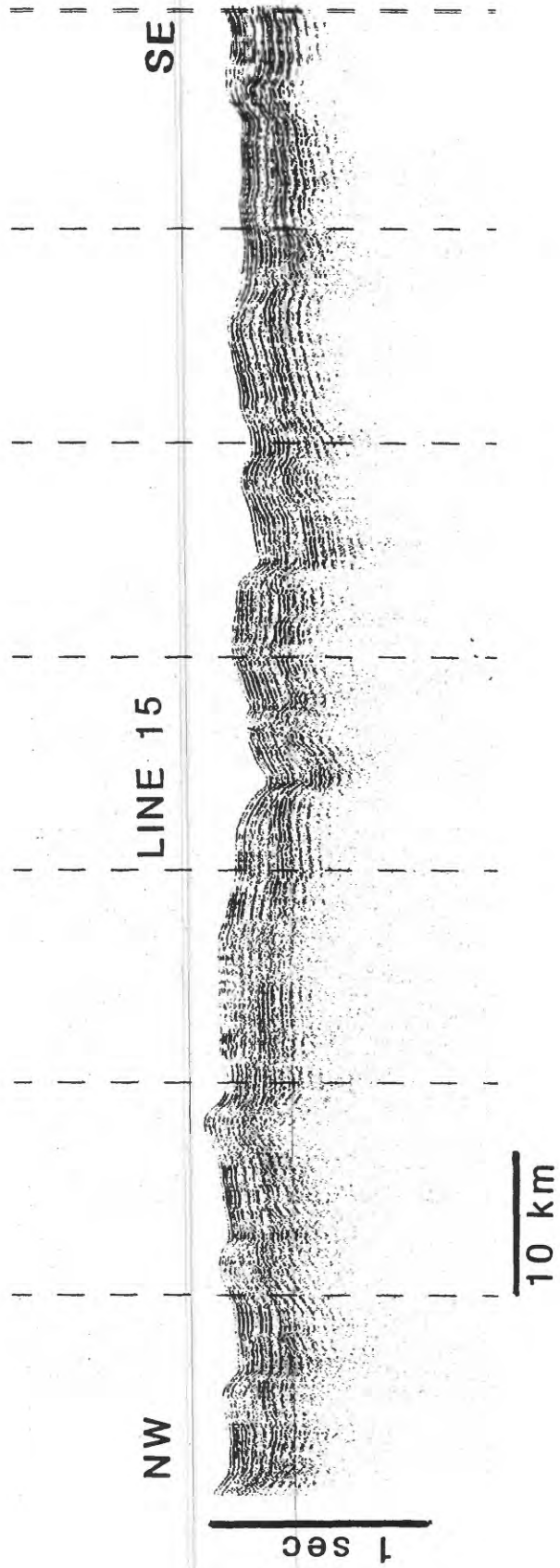


Figure 4



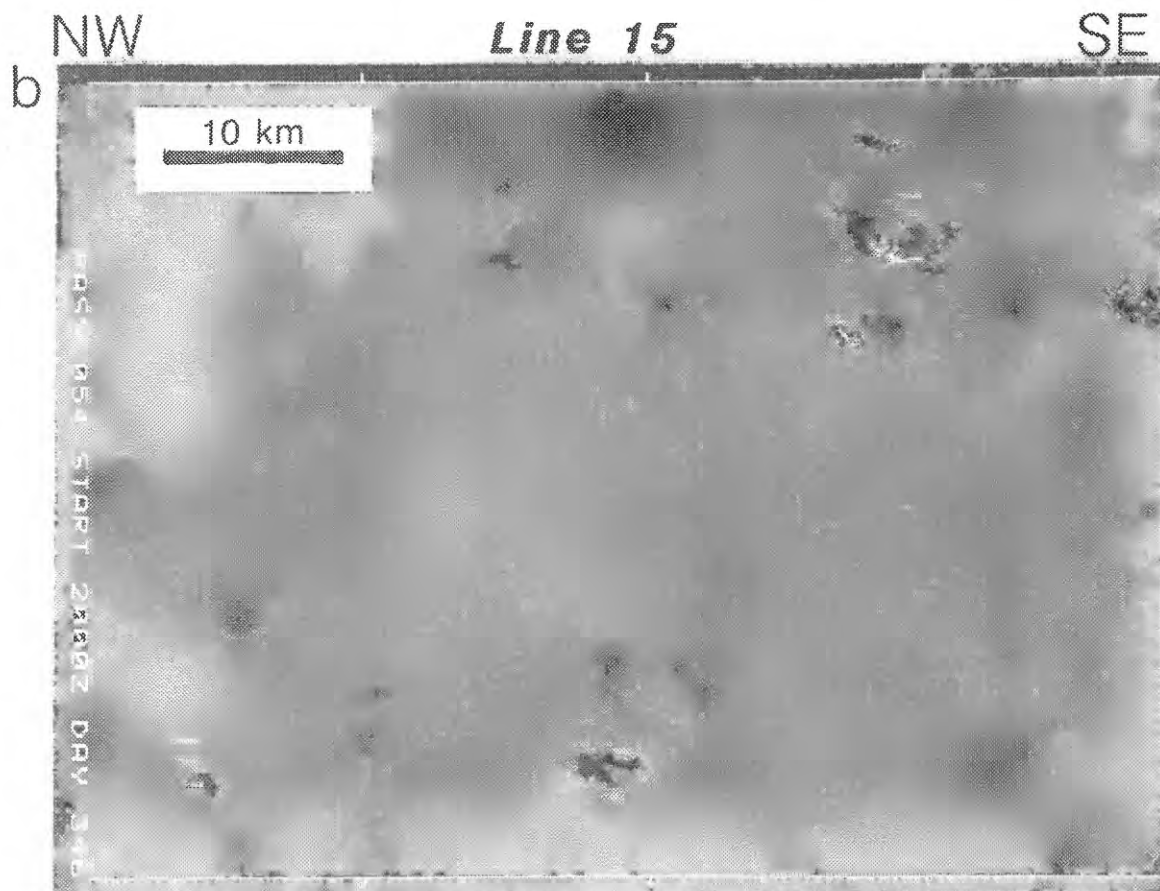
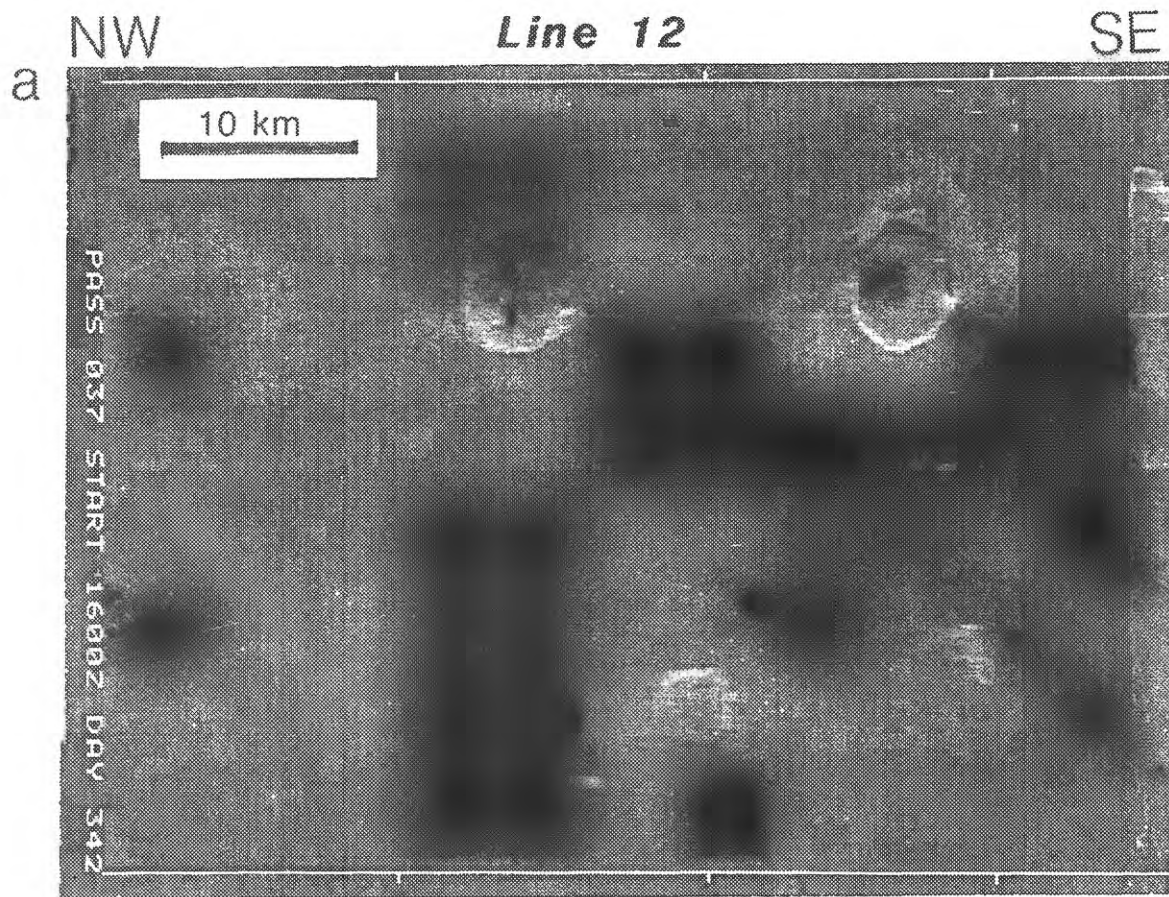


Figure 5

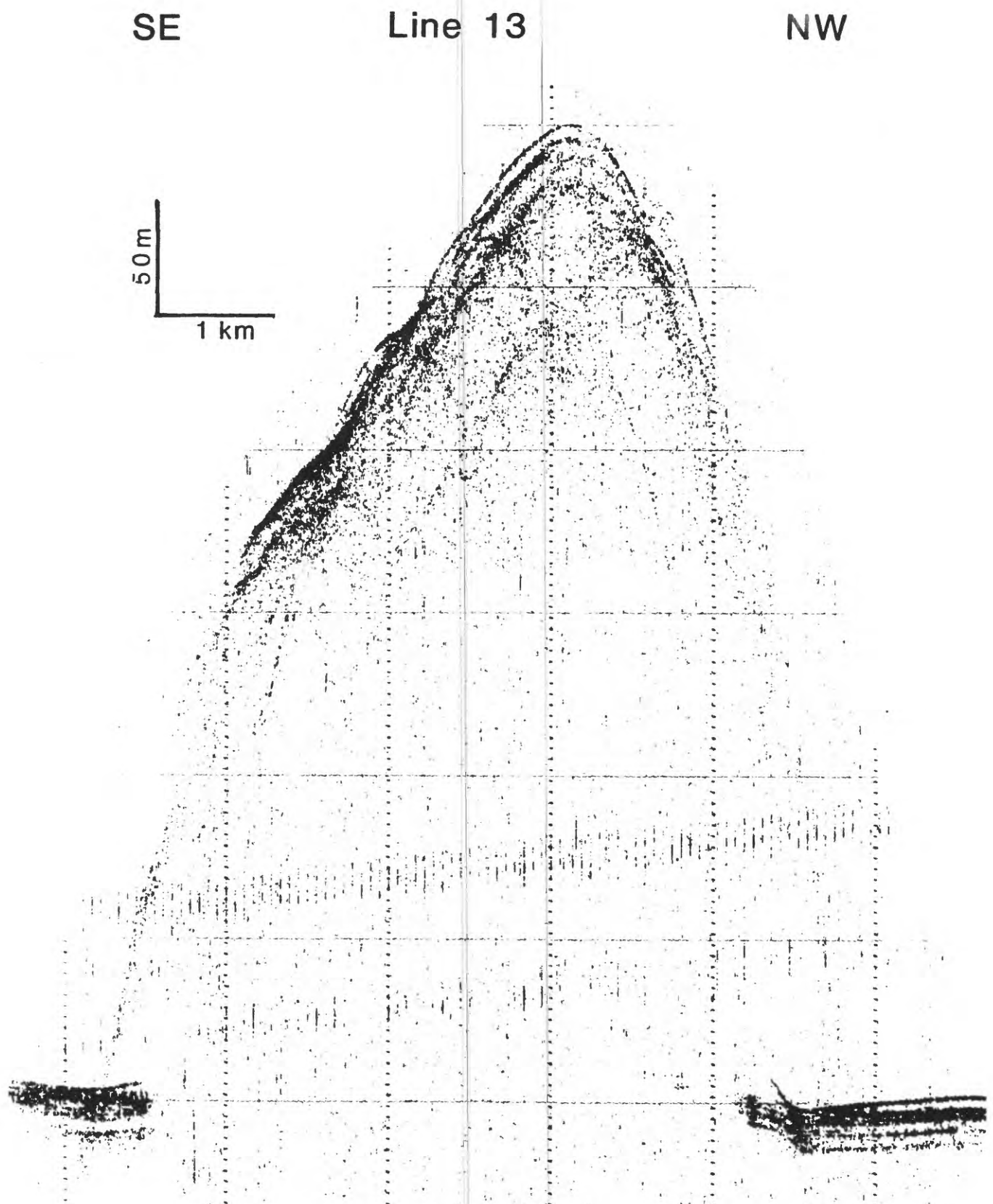


Figure 6



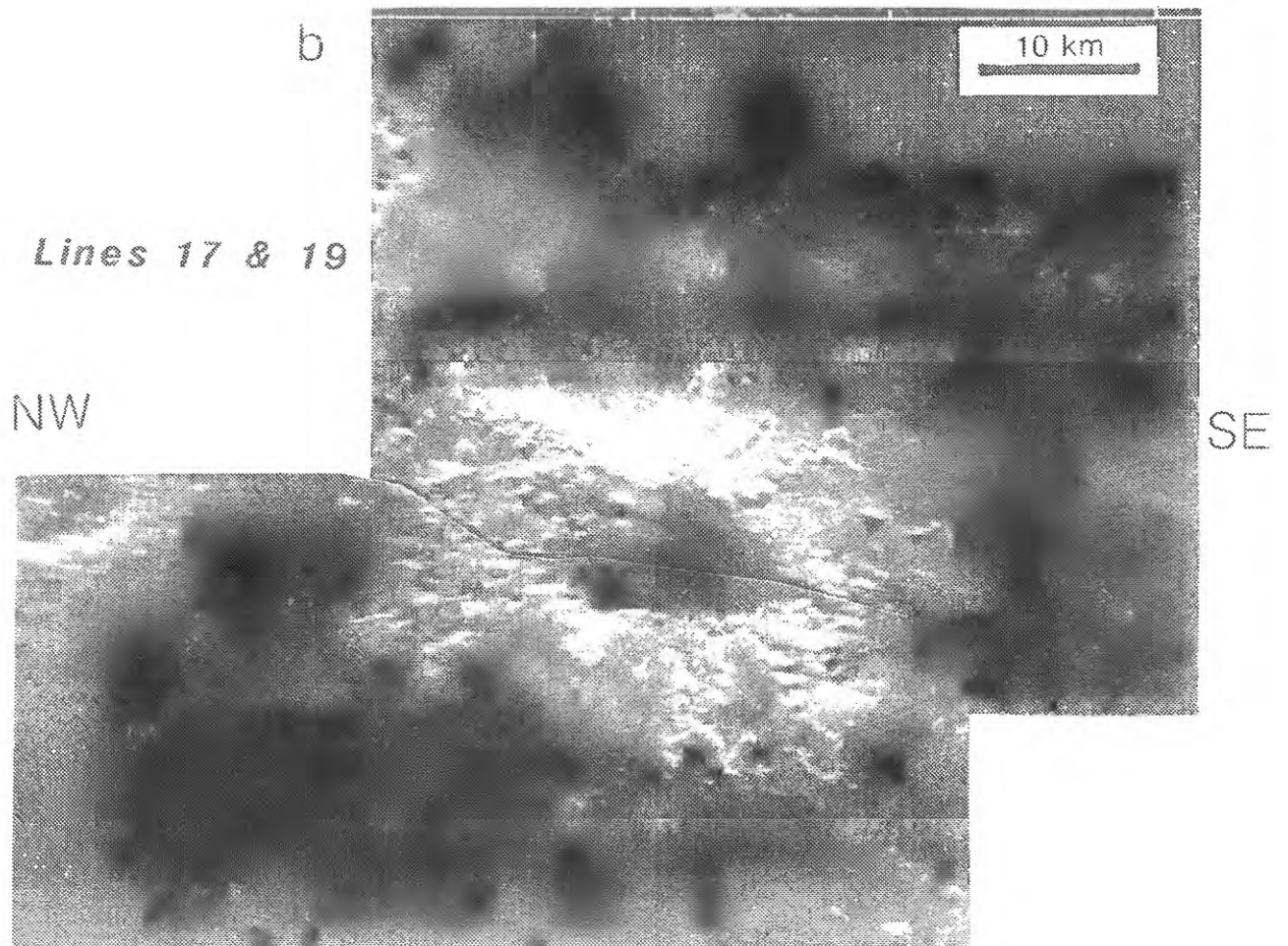
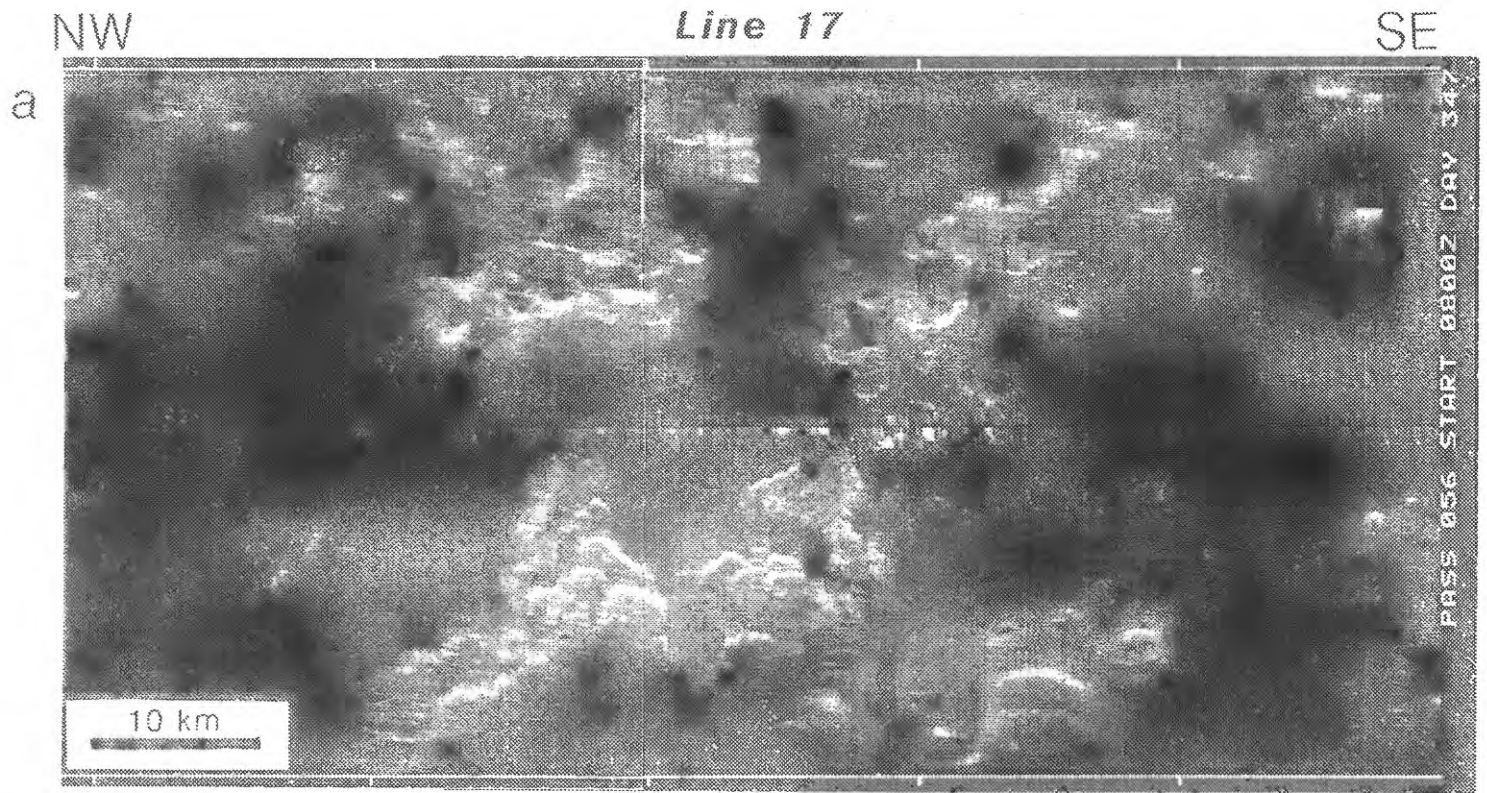


Figure 7

## Line 15

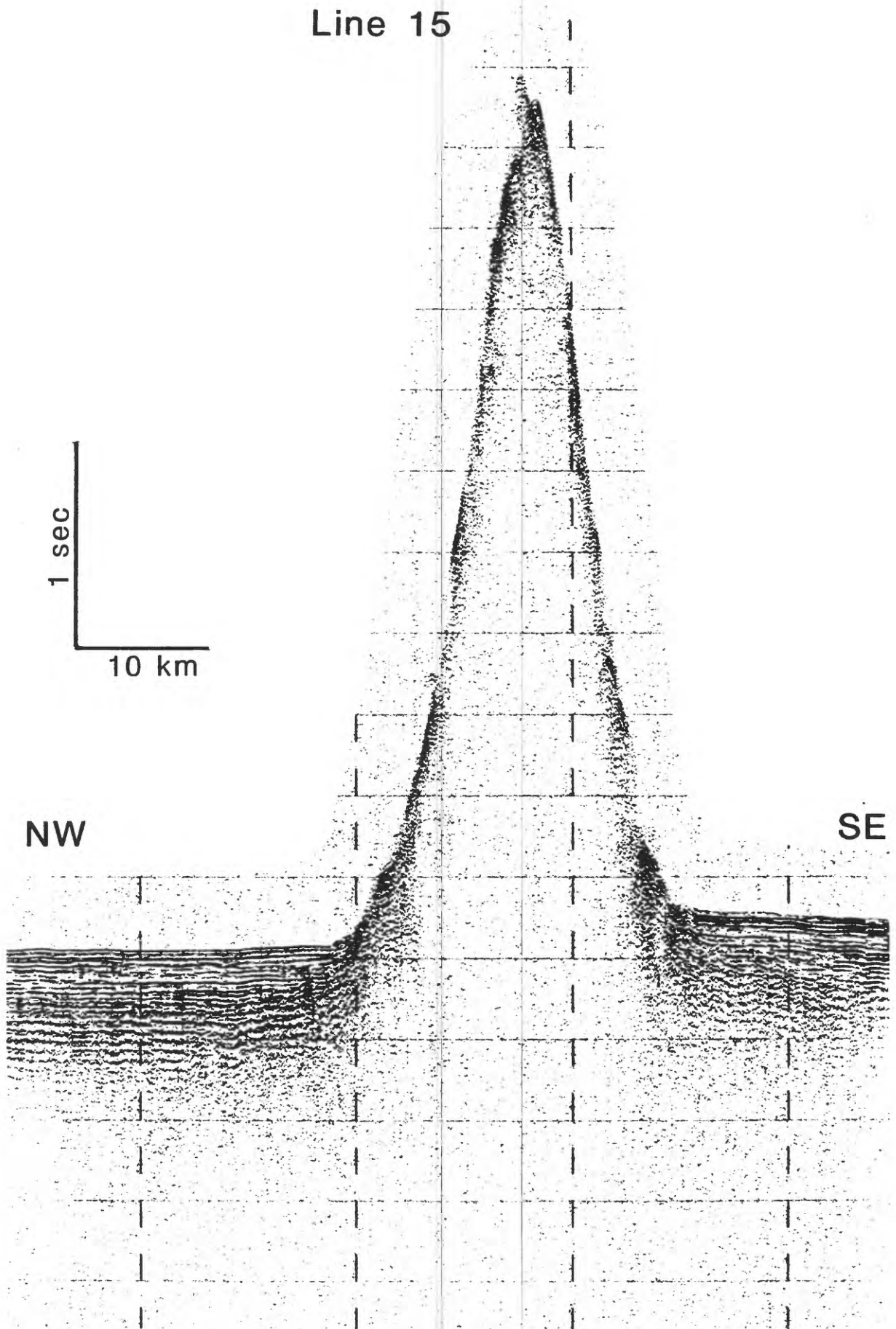


Figure 8

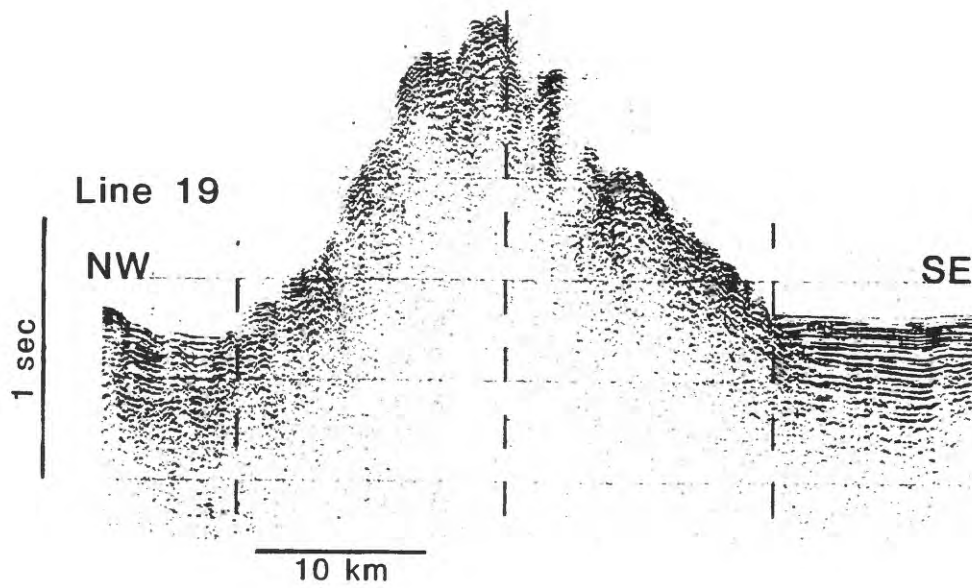
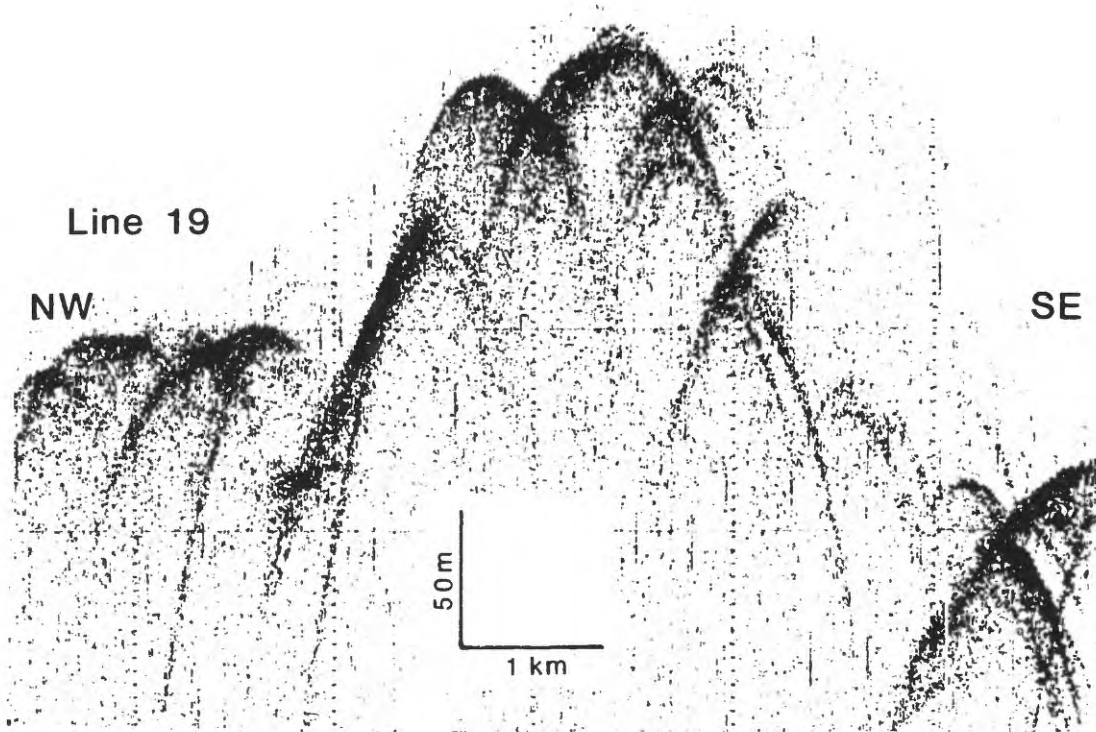


Figure 9

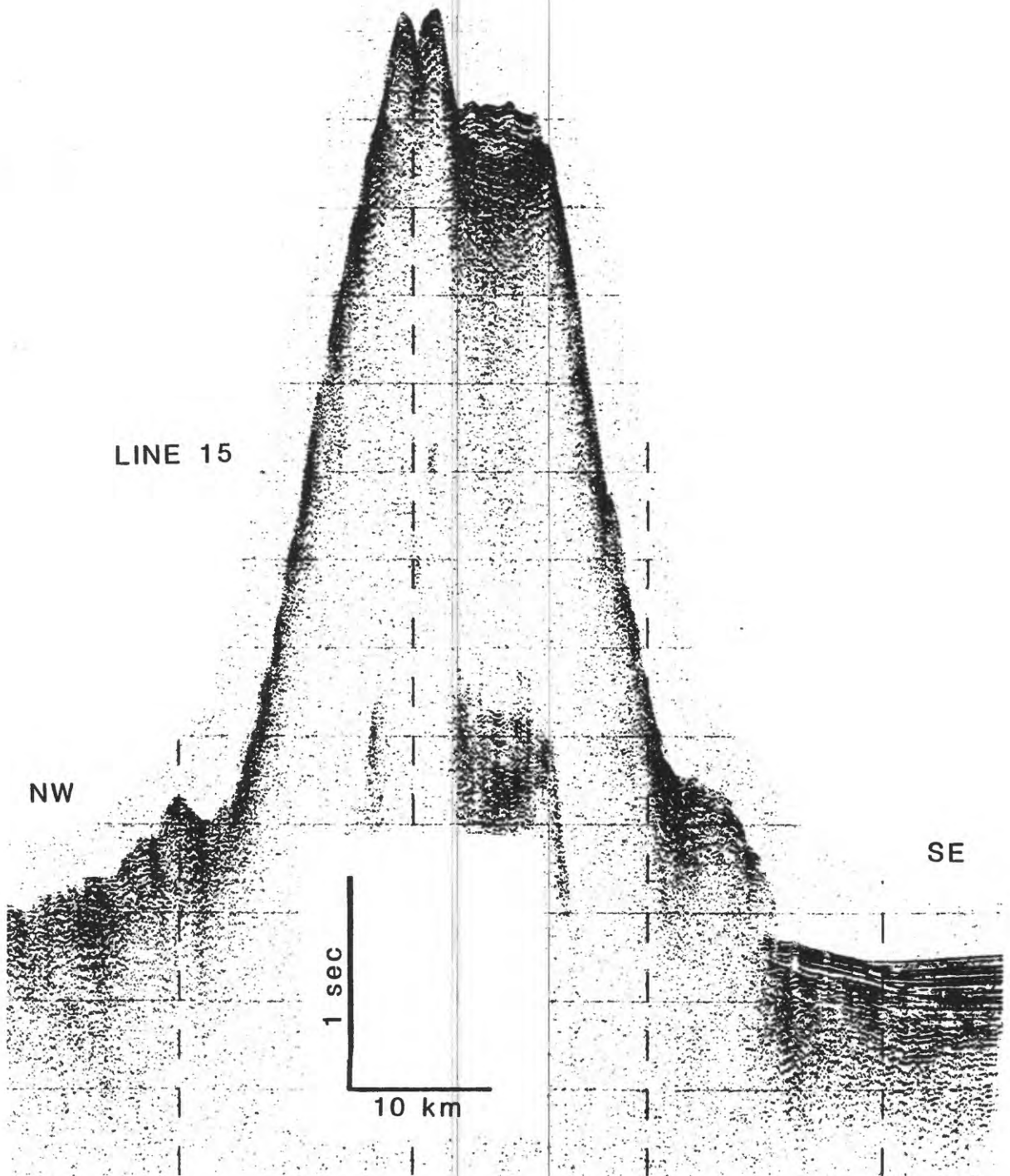


Figure 10

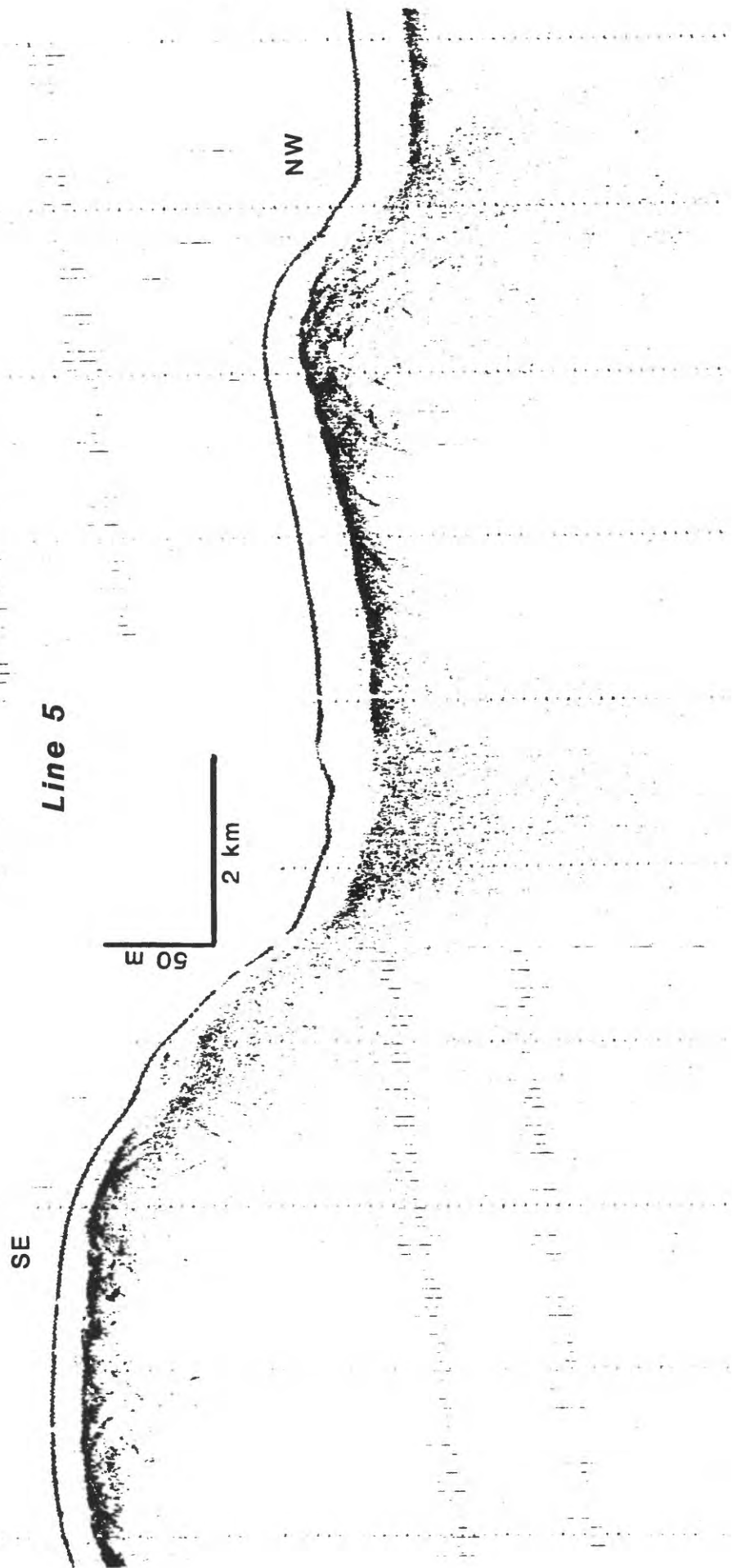


Figure 11



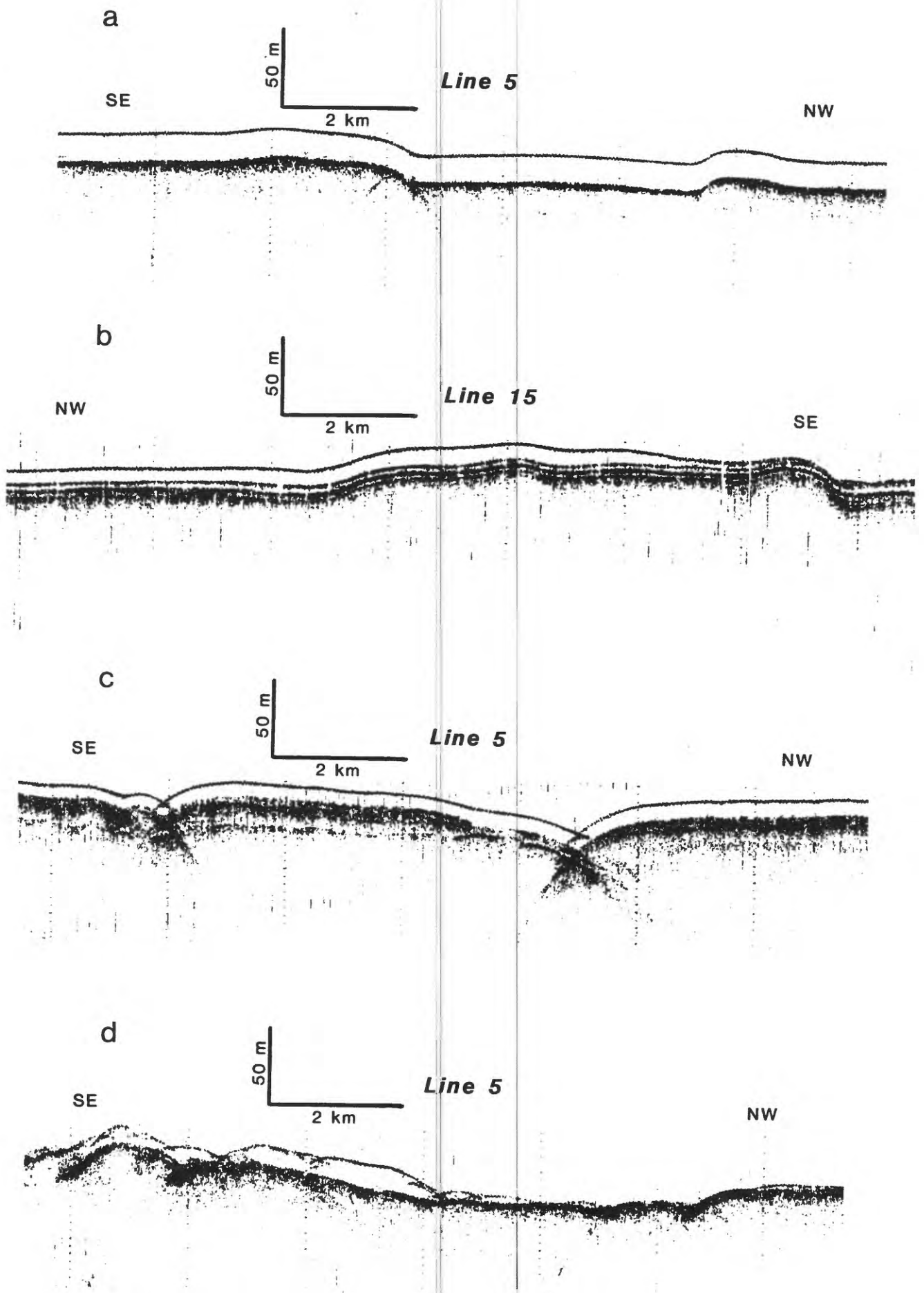


Figure 12

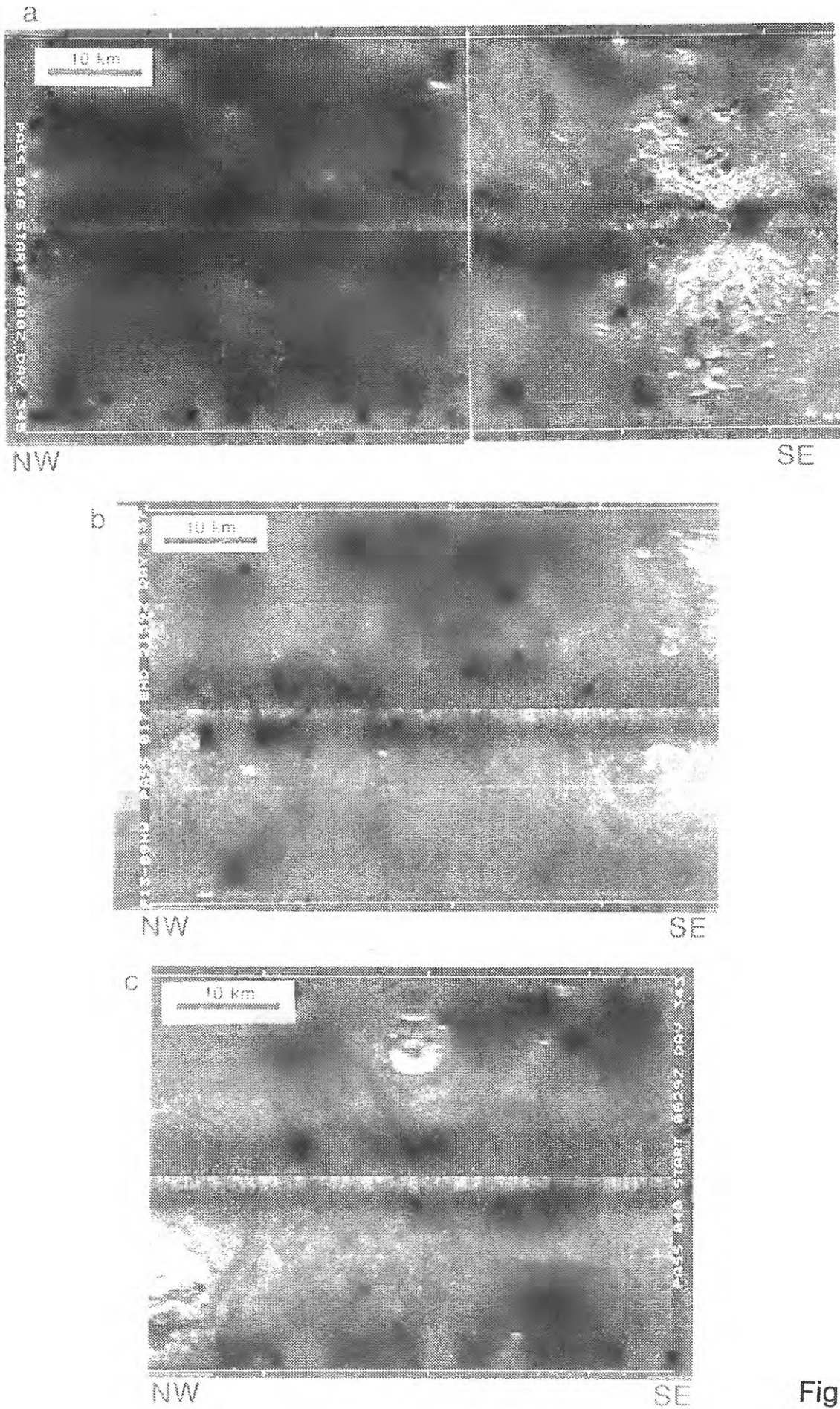


Figure 13

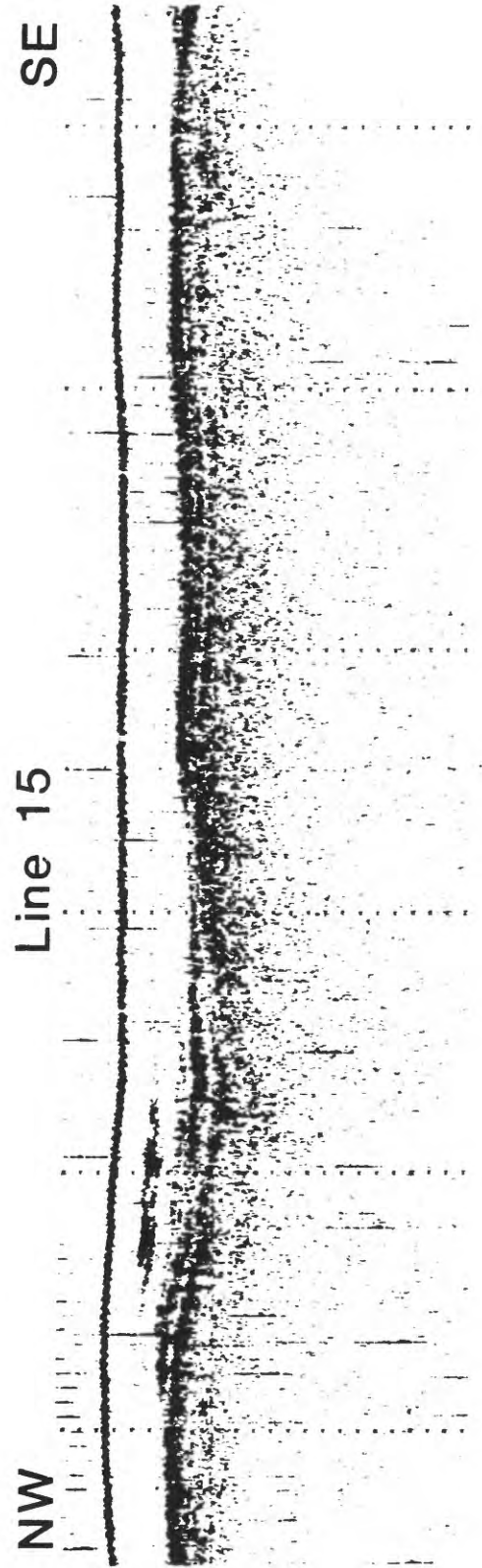
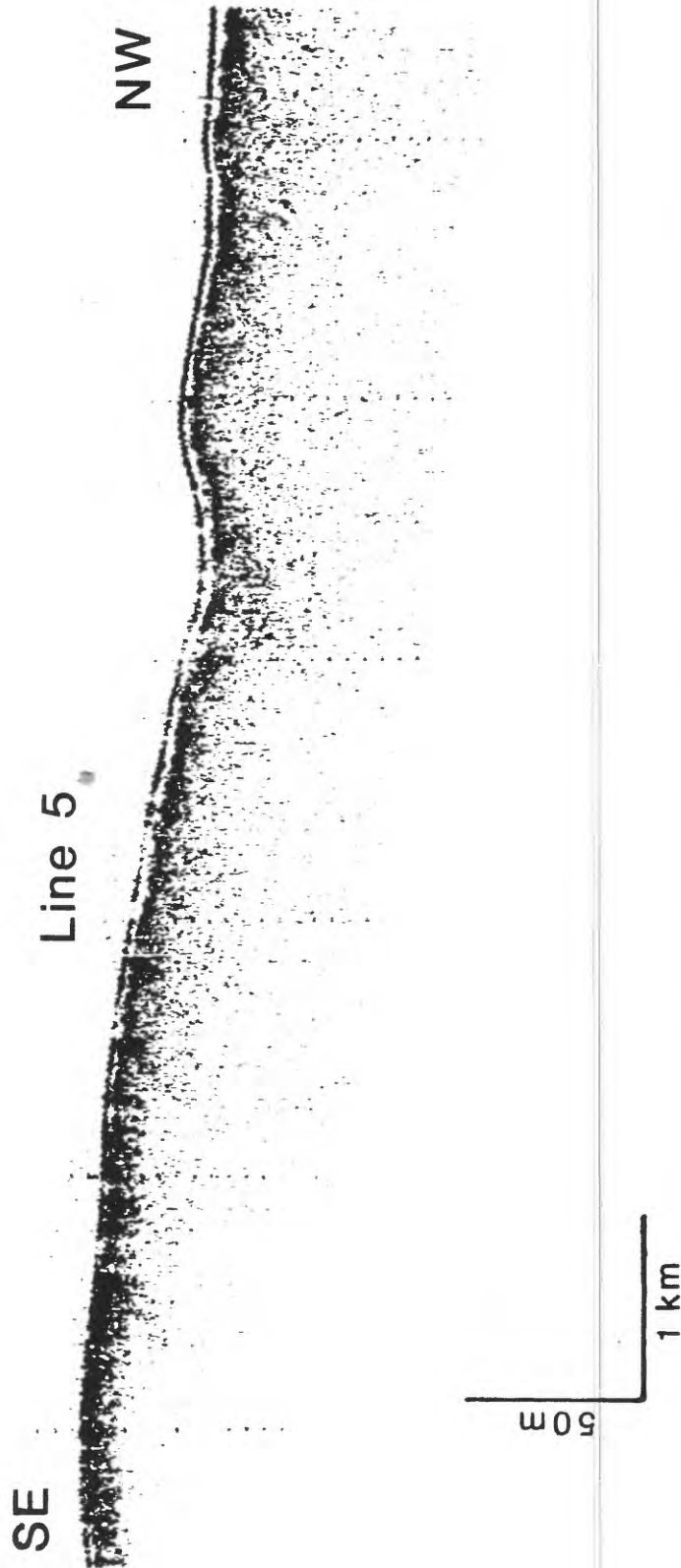


Figure 14 .