

**U.S. DEPARTMENT OF THE INTERIOR**  
**U.S. GEOLOGICAL SURVEY**

**Measured Section of the Pennsylvanian Hermosa Group Near Ouray, Colorado**

**by**

**Karen J. Franczyk<sup>1</sup>**

**Open-File Report 93-712**

**This report is preliminary and has not been reviewed for conformity with U.S. Geological Survey editorial standards (or with the North American Stratigraphic Code). Any use of trade, product, or firm names is for descriptive purposes only and does not imply endorsement by the U.S. Government.**

**<sup>1</sup>USGS, Denver, CO**

## Introduction

A comprehensive study of the Paradox Basin, located in southeastern Utah and southwestern Colorado, is being conducted from 1989-1994 under the U.S. Geological Survey's Evolution of Sedimentary Basins Program. As part of the investigation of the Pennsylvanian System in the Paradox Basin, sections were measured in detail at various locations along the Pennsylvanian outcrop belt in southwestern Colorado. These sections were measured to establish better correlations between outcrops of Pennsylvanian strata in the area that formed the active, eastern, basin margin of the Pennsylvanian Paradox Basin; to establish better correlations from the eastern margin to the center of the basin; and to better define depositional processes on the basin's active margin. This report shows the measured section of the Pennsylvanian Molas Formation and Hermosa Group at Ouray, Colorado (fig. 1). Open-File Report 92-689 (Franczyk, 1992) contains the detailed measured section of the Hermosa Group at Hermosa Mountain, which is about 20 km (12 mi) north of Durango, Colorado (fig. 1).

The Pennsylvanian rocks in the Ouray, Colorado, area are the most northern exposures, and the most proximal to the Uncompahgre uplift, of Pennsylvanian outcrops along the area that formed the southeastern margin of the Pennsylvanian Paradox Basin. South of Ouray erosion has removed Paleozoic rocks for a distance of 32 km (20 mi). Pennsylvanian rocks crop out again south of Silverton, Colorado, and form a continuous outcrop belt southward to the Hermosa Mountain area. Extensive vegetation cover, talus deposits, and local glacial deposits prohibit detailed correlations and resolution of lithofacies geometries along this outcrop belt.

There are no published measured sections of Pennsylvanian strata at Ouray. However, Fetzner (1960) and Wengerd and Strickland (1954) used their unpublished measured sections at Ouray as data points in isopach maps and cross sections; both workers recorded about 305 m (1000 ft) of Pennsylvanian rock. Luedke and Burbank (1962) mapped the Ouray Quadrangle and, based on structural cross sections, estimated 442 m (1450 ft) of Pennsylvanian strata. The measured section of this report shows 305 m (1000 ft) inclusive of both the Pennsylvanian Molas Formation and Hermosa Group. The Triassic Dolores Formation unconformably overlies the Hermosa Group at the measured section locality. This unconformity rises stratigraphically to the north, and an additional 60-120 m (200-400 ft) of Hermosa Group may be present north of the measured section location.

This report presents the lithologic descriptions, macroscopic observations, and interpretations of depositional environments from the Ouray measured section. The format used here enables more detail to be shown than would be possible in a standard manuscript figure or appendix. Ongoing studies of clastic and carbonate petrography and of brachiopod assemblages from this measured section will be integrated with similar outcrop data from other localities and compiled in future reports.

## Methods

The Ouray measured section (fig.1), located on the Ouray 7.5 minute quadrangle map (Ouray County, Colorado), is comprised of two segments (fig. 2). The lower segment, named the Ouray segment, begins in section 31, T. 44 N., R. 7 W. (projected location in unsurveyed area) and extends into section 36, T. 44 N., R. 7 W. The upper segment, named the Cascade Falls segment, begins in section 31 and extends into section 32, T. 44 N., R. 7 W. (projected location in unsurveyed area). The lower segment begins at the top of the Mississippian Leadville Limestone and ends at 119 m (391 ft) at the top of a thick sandstone unit above which is extensive cover. The upper segment begins with the lowest exposed strata at Cascade Falls and follows the southeast-trending contact between the talus and cliffs up through the Hermosa Group. The basal 59 m (195 ft) of the Cascade Falls segment duplicates the upper part of the Ouray segment. The duplicated section is included

to show the lithologic changes that occur over the 1.6 km (1 mi) that separate these two segments.

Along the line of section, an angular unconformity separates the Hermosa Group from the Triassic Dolores Formation. This unconformity rises stratigraphically to the north where it separates the Permian Cutler Formation from the overlying Dolores. Shear, inaccessible cliffs prohibit examination of the Hermosa-Cutler contact where it is exposed. From a distance, this contact looks conformable and gradational.

The section was measured with a Jacobs staff and Abney level; in high, cliff-forming exposures a tape was dropped from the top of the cliff. There is excellent exposure in both the Ouray and Cascade Falls segments resulting in documentation of both the finer-grained units and the contacts between units. Field descriptions were recorded at a scale of 1 in. to 5 ft. or 1 in. to 10 ft (1 cm to .6 m or 1 cm to 1.2 m). These descriptions were simplified to the 1 in. to 20 ft (1 cm to 2.4 m) scale presented in this report.

Both sandstone and carbonate beds were sampled and thin sections prepared for petrographic analysis. Petrographic analysis in the clastic units will document changes in mineralogy through the section and in the carbonate units will determine faunal assemblages and abundances and diagenetic history. The faunal types shown in the accessories column of figures 4 and 5 are based on macroscopic field observations and will be expanded with the final analysis of these carbonate samples. Samples of dark-gray claystone, mudstone, and shale throughout the section were also collected for organic-geochemistry analysis to determine total organic content and thermal maturation levels.

Figure 3 provides the explanation for the measured section diagrams, and figures 4 (Ouray segment) and 5 (Cascade Falls segment) show the measured section data. Part or all of the unit at the top of each page on figures 4 and 5 is repeated at the base of the succeeding page to facilitate splicing the section into a continuous vertical column. The profile on the left side of the lithology column shows the grain size for clastic units. The lithologic profile for carbonate units is arbitrarily assigned to the medium-grained interval; the texture of the carbonate unit is shown in the carbonate fabric column. An interpretation of depositional environments for units or packages of units is listed on the far left side of figures 4 and 5. Based on the interpreted environments, a relative sea level curve was constructed. Marine carbonate and marine clastic units are considered to represent the highest relative sea level and alluvial units the lowest. Abrupt changes from high or intermediate relative sea level to low correspond to a period of base level drop and subaerial exposure.

### Geologic notes

Pennsylvanian strata in the Ouray area can be separated into two units: the Molas Formation and the overlying Hermosa Group. The Molas, 18 m (60 ft) thick, unconformably overlies the Mississippian Leadville Limestone. Red sandstone to siltstone matrix surrounding karstified limestone characterizes the lowest part of the Molas. This lithology is replaced upward by red conglomerate and pebbly sandstone beds. A color change from dominantly red to dominantly olive gray and a slight decrease in grain size marks the gradational contact between the alluvial deposits of the Molas Formation and those of the overlying Hermosa Group. The Pinkerton Trail Formation, a marine carbonate unit at the base of the Hermosa Group throughout the Paradox basin, is absent (possibly due to nondeposition) in the Ouray area. The first marine carbonate unit overlies 30 m (100 ft) of alluvial deposits in the Hermosa Group.

Luedke and Burbank (1962) informally separated the Hermosa Group at Ouray into three parts based on lithologic characteristics. Their lower part is about 137 m (450 ft) thick and consists of gray, green, and red calcareous sandstone, siltstone, and shale interbedded with gray fossiliferous limestone. The middle part is about 213 m (700 ft) thick and is characterized by pink to red, generally cliff-forming, thick (15-24 m (50-80 ft))

sandstone units interbedded with thin, red to gray limestone and fossiliferous shale units. The upper part is about 91 m (300 ft) thick and contains red sandstone and conglomerate beds that are thinner than those in the underlying middle part. These coarse-grained clastic beds are interbedded with fossiliferous gray limestone and gray to red shale units. Luedke and Burbank (1962) interpreted a gradational and conformable contact between the Hermosa and the Cutler; they arbitrarily chose the contact between these two units to be 40 m (130 ft) above the uppermost fossiliferous zone within the Hermosa.

This measured section shows a similar three-fold subdivision but slight differences in thickness. The lower part, 104 m (340 ft) thick, contains overall thinner lithologic units than in the overlying parts. Beds of sandstone, siltstone, shale, and limestone ranging from less than 3 m to 9 m (10 ft to 30 ft) thick characterize the lower part. The middle part, which is 154 m (505 ft) thick, contains cliff-forming channel-sandstone units that are locally up to 27 m (90 ft) and generally greater than 12 m (40 ft) thick. Fossiliferous, subtidal limestone units ranging from 0.6-3 m (2-10 ft) thick overlie these channel sandstones above a sharp contact. Only 27 m (90 ft) of the upper part was measured along the line of section because of erosion below the angular unconformity at the base of the Triassic. Thinner units, more fine-grained clastic intervals, and less well developed limestone beds differentiate it from the underlying middle part. Lithologically, with the exception of limestone in the Hermosa, the upper part of the Hermosa is identical to the overlying basal part of the Cutler where the Cutler has not been removed below the unconformity surface (Luedke and Burbank, 1962).

Luedke and Burbank (1962) did not apply the name Rico Formation in the Ouray area to the interval that contains thin, poorly developed limestone beds at the top of the Hermosa Group. They did note that the limestone beds in the upper part of the Hermosa contain a greater ratio of molluscs to brachiopods than in the underlying parts. Cross and Spencer (1900) used this increase in mollusc abundance to differentiate the limestones in the Rico Formation from those in the underlying Hermosa in the Rico, Colorado, area. Based on the greater abundance of molluscs in upper part of the Hermosa at Ouray, Luedke and Burbank (1962) suggested that this part of the Hermosa "may be an approximate equivalent of the Rico Formation" at Rico, Colorado. Cross and Spencer (1900, p. 65), citing personal communication from Girty, reported fossils from the Rico Formation having "both Permian and Carboniferous affinities, but if this unit must be assigned to only one of these ages, it would be placed in the Carboniferous (Pennsylvanian)". Blodgett (1984, p. 68) stated that upper Desmoinesian fusulinids were found in the uppermost limestone beds in the Rico area, indicating the Hermosa/Rico section is no younger than Pennsylvanian.

Ongoing paleontologic study of the faunal assemblages at Ouray indicates the Hermosa Group is no younger than Pennsylvanian, and probably no younger than the DesMoinesian stage of the Pennsylvanian (Germaine Clark, USGS, 1993, oral communication). The apparent conformable and gradational contact with the overlying Cutler Formation suggests that the Cutler is of both Pennsylvanian and Permian age.

Correlations of the Hermosa Group between Ouray, Silverton, and Hermosa Mountain present interesting questions. Figure 6 shows the great increase in thickness in the Hermosa Group between the Ouray and Silverton areas. Additionally, the gross, regionally extensive lithologic packages that can be correlated from Hermosa Mountain to the Silverton area (fig. 6) can not be identified in the Ouray area. Because the Pennsylvanian rocks south of Ouray area are eroded, the cause of these changes is uncertain. A stratigraphic rise in the Hermosa-Cutler contact from Ouray to Silverton is possible, or there may be an unconformity spanning a large part of the DesMoinesian within the Hermosa Group at Ouray. The interpretation shown in figure 6 is that numerous unconformities within the Hermosa at Ouray and a slight stratigraphic rise in the Hermosa-Cutler contact account for the observed thickness differences. The sharp, scour contacts that locally show a few meters of relief at the base of thick channel sandstone units in the

middle part of the Hermosa Group at Ouray are probably the sites of these unconformities. Ouray's close proximity to the Uncompahgre uplift may have made this area more sensitive to tectonically induced base-level changes that resulted in long periods of erosion and nondeposition. Ongoing paleontologic analysis of the faunal assemblages may provide more precise chronostratigraphic correlations between Ouray and areas to the south.

Trough cross-stratification within deltaic marine and nonmarine fluvial sandstones throughout the Hermosa Group in the Ouray area yield radial (southwest to north) paleocurrent indicators that produce an average mean vector of  $281^{\circ}$  (fig. 7). This westerly direction differs from the southwesterly orientation ( $230^{\circ}$ ) obtained from the Hermosa Group in both the Molas Lake (Spoelhof, 1974) and the Hermosa Mountain areas (Franczyk, 1992).

#### Acknowledgements

The long, arduous, and often tedious job of section measuring was ably and good-naturedly assisted by George Wigman and Doug Brew.

#### References Cited

- Blodgett, R.H., compiler, 1984, Road log from Durango to Telluride, Colorado, via Mancos, Dolores, and Rico, Colorado, in, Brew, D.C., ed., Field Trip Guidebook: Four Corners Geological Society, 37th Annual Meeting Rocky Mountain Section Geological Society of America.
- Cross, Whitman, and Spencer, A.C., 1900, Geology of the Rico Mountains, Colorado: U.S. Geological Survey 21st Annual Report, Pt. 2, p. 7-165.
- Franczyk, K.J., 1992, Measured Section of the Pennsylvanian Hermosa Group Near Hermosa, Colorado: U.S. Geological Survey Open-File Report 92-689.
- Fetzner, R.W., 1960, Pennsylvanian paleotectonics of Colorado Plateau: American Association of Petroleum Geologists Bulletin, v. 44, no. 8, p. 1371-1413.
- Franczyk, K.J., Pitman, J.P., and Brew, D.C., in press, Sedimentology, Mineralogy, and Age of the Hermosa Group at Hermosa Mountain, Southwestern Colorado: U.S. Geological Survey Circular.
- Luedke, R.G., and Burbank, W.S., 1962, Geology of the Ouray Quadrangle Colorado: U.S. Geological Survey Geologic Quadrangle Map GQ-152, scale 1:24,000.
- Spoelhof, R.W., 1974, Pennsylvanian stratigraphy and tectonics in the Lime Creek-Molas Lake area, San Juan County, Colorado: Golden, Colorado School of Mines, PhD Thesis, 193 p.
- Wengerd, S.A., and Strickland, J.W., 1954, Pennsylvanian stratigraphy of Paradox salt basin, Four Corners region, Colorado and Utah: American Association of Petroleum Geologists Bulletin, v. 38, no. 10, p. 2157-2199.

## Figure Captions

Figure 1. Location map showing the Paradox Basin, shown here by the solid, heavy line that corresponds to the maximum extent of salt in the Paradox Formation, and surrounding geologic features. The star shows the location of the Ouray measured section. Line of section shows location of stratigraphic cross section illustrated in figure 6.

Figure 2. Location map of the Ouray measured section. The Ouray segment includes the Molas Formation and lowest part of the Hermosa Group. The Cascade Falls segment includes the remainder of the Hermosa Group. At the top of the measured section, the Triassic Dolores Formation unconformably overlies the Hermosa. Base map from Luedke and Burbank's (1962) geologic map of the Ouray 7.5 minute quadrangle map.

Figure 3. Explanation for figures 4 and 5.

Figure 4. Ouray segment of the Ouray measured section. Part or all of the unit at the top of each page is repeated at the base of the succeeding page to facilitate splicing the section into a continuous vertical column.

Figure 5. Cascade Falls segment of the Ouray measured section. The basal 59 m (195 ft) of the Cascade Falls segment duplicates the upper part of the Ouray segment. Part or all of the unit at the top of each page is repeated at the base of the succeeding page to facilitate splicing the section into a continuous vertical column.

Figure 6. Stratigraphic correlation diagram of the Molas Formation ( $\mathbb{P}m$ ), Pinkerton Trail Formation ( $\mathbb{P}pt$ ), and Hermosa Group ( $\mathbb{P}h$ ) between Ouray to the north and Hermosa Mountain to the south. Line of section is shown on figure 1. Spoelhof (1974) refers to the Hermosa Group above the Pinkerton Trail Formation as the Honaker Trail Formation ( $\mathbb{P}ht$ ). The middle member of the Honaker Trail of Spoelhof is a carbonate-rich interval that can be traced from the Silverton area southward to Hermosa Mountain; this interval can not be identified in the Ouray area. Ml-Mississippian Leadville Limestone; PC-Precambrian; Td-Triassic Dolores Formation;  $\mathbb{P}Pc$ -Pennsylvanian/Permian Cutler Formation. Equivalent (eq.) members of the Paradox Formation of the Hermosa Group are shown in column on right: AG-Alkali Gulch; BC-Barker Creek; A-Akah; DC-Desert Creek; I-Ismay.

Figure 7. Rose diagram showing all paleocurrent measurements taken from the Ouray and Cascade Falls segments. Average mean vector direction is  $281^\circ$ ; N is the number of readings.

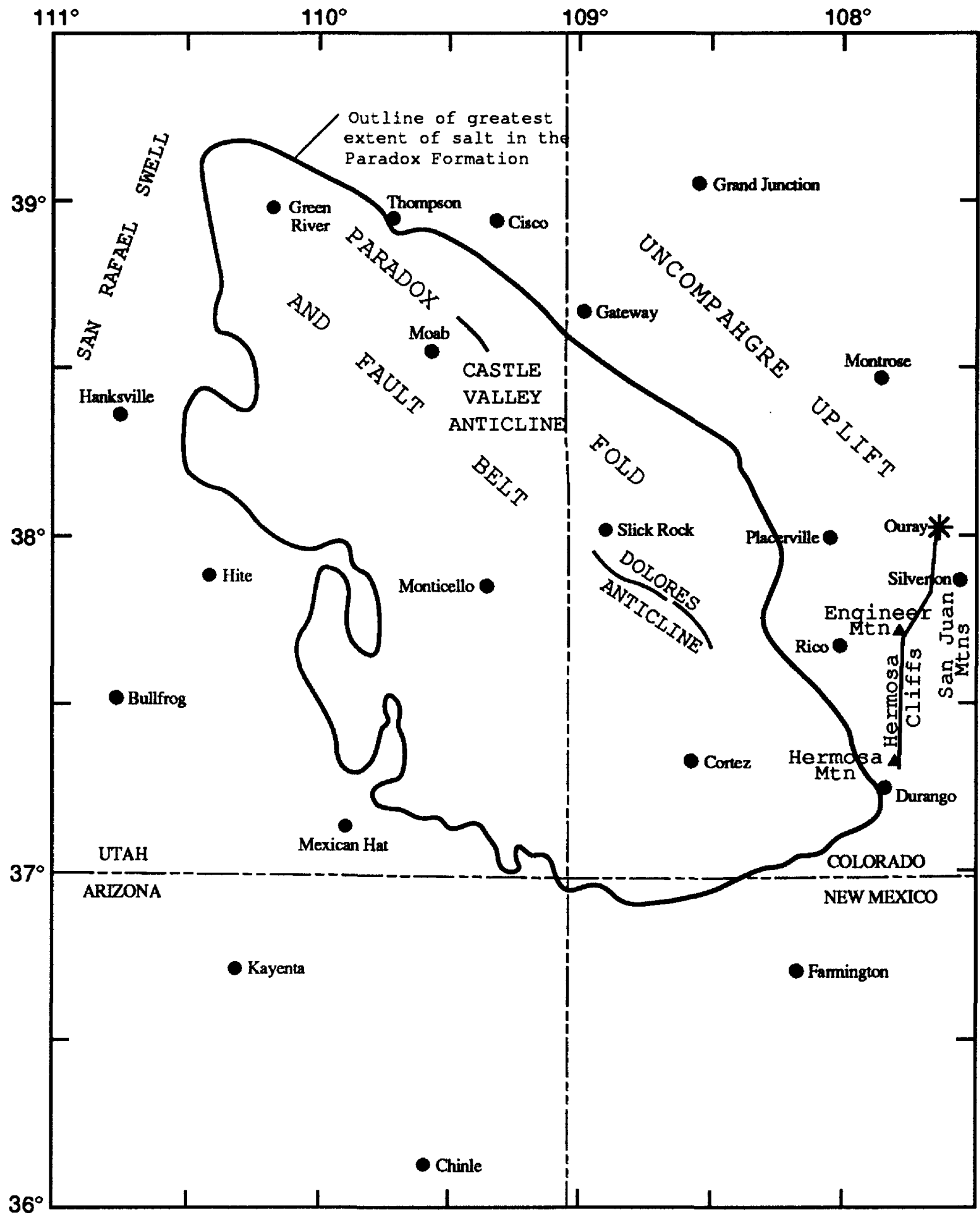
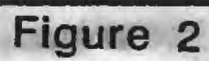


Figure 1













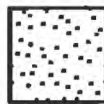































































EXPLANATION

LITHOLOGY

ACCESSORY FEATURES

	Conglomerate		Claystone/shale		root cast		flute casts		bryozoan
	Pebbly sandstone		Limestone		vertical burrow		plant fragment		indeterminate fossil fragment
	Sandstone		Shaley limestone		horizontal burrow		wood		echinoids
	Calcareous sandstone		Breccia		carbonaceous debris		coral		chert
	Siltstone		Covered		convolute bedding		fusulinid		carbonate nodule
	Interbedded siltstone and sandstone		Shaley limestone		intraformational clast		pelecypod		feldspar
	Calcareous sandstone		Breccia		bioturbation		gastropod		mica
	Siltstone		Covered		flame structures		brachiopod		pyrite
	Interbedded siltstone and sandstone		Covered		load casts		Chaetetes		abundant
	Siltstone		Covered		current ripples		algae		same as below
	Interbedded siltstone and sandstone		Covered		wave ripples		tabular planar cross-stratification		low-angle stratification
	Interbedded siltstone and sandstone		Covered		ripple-type unknown		parallel discontinuous laminations		planar stratification
	Interbedded siltstone and sandstone		Covered		climbing ripples		hummocky cross-stratification		scour and fill
	Interbedded siltstone and sandstone		Covered		wavy nonparallel laminations		trough cross-stratification		lateral accretion bedding

SEDIMENTARY STRUCTURES

Figure 3. p. 1

Figure 3, p. 1

EXPLANATION				
DEPOSITIONAL ENVIRONMENT	CARBONATE FABRIC	BEDDING THICKNESS	CONTACT TYPE	
NM Nonmarine	M Lime mudstone	VTH Very thin	S	Sharp
AL Alluvial	W Wackestone	TH Thin	SI	Sharp, irregular
DP Delta plain	P Packstone	M Medium	G	Gradational
DF Delta front	G Grainstone	TK Thick	C	Covered
SM Shallow marine	X Crystalline	VTK Very thick		
		NA Not apparent		
M Marine				
COLOR			CONTACT TYPE	
B Brown	Or Orange	D Dark	C	Calcareous
Bl Black	Pk Pink	L Light	S	Siliceous
G Gray	R Red	M Medium	SI	Slightly
Gr Green	Y Yellow	P Pale	V	Very
O Olive			N	Non

Figure 3, p. 2





Fig. 4  
Ouray p. 2

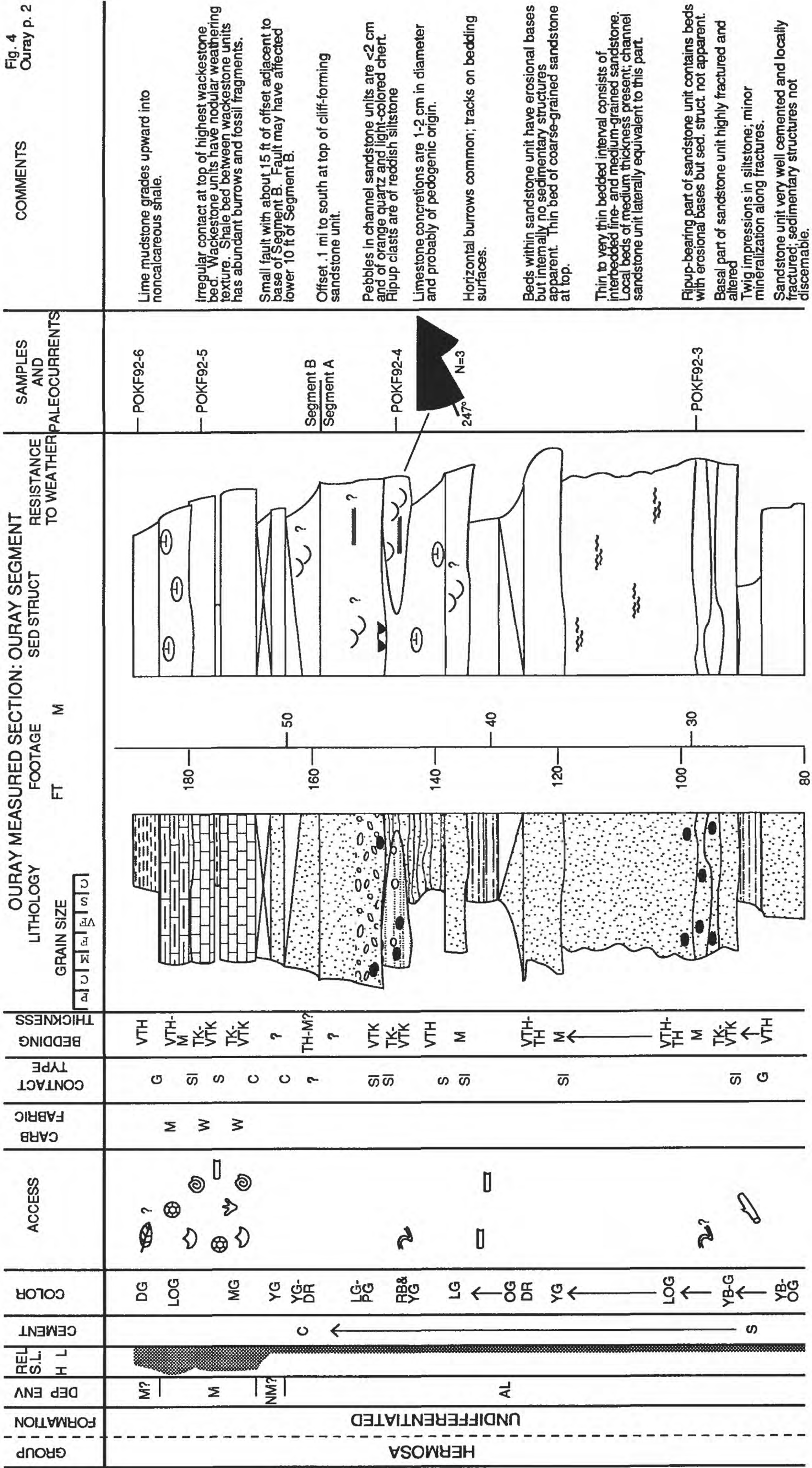


Fig. 4  
Ouray p.3

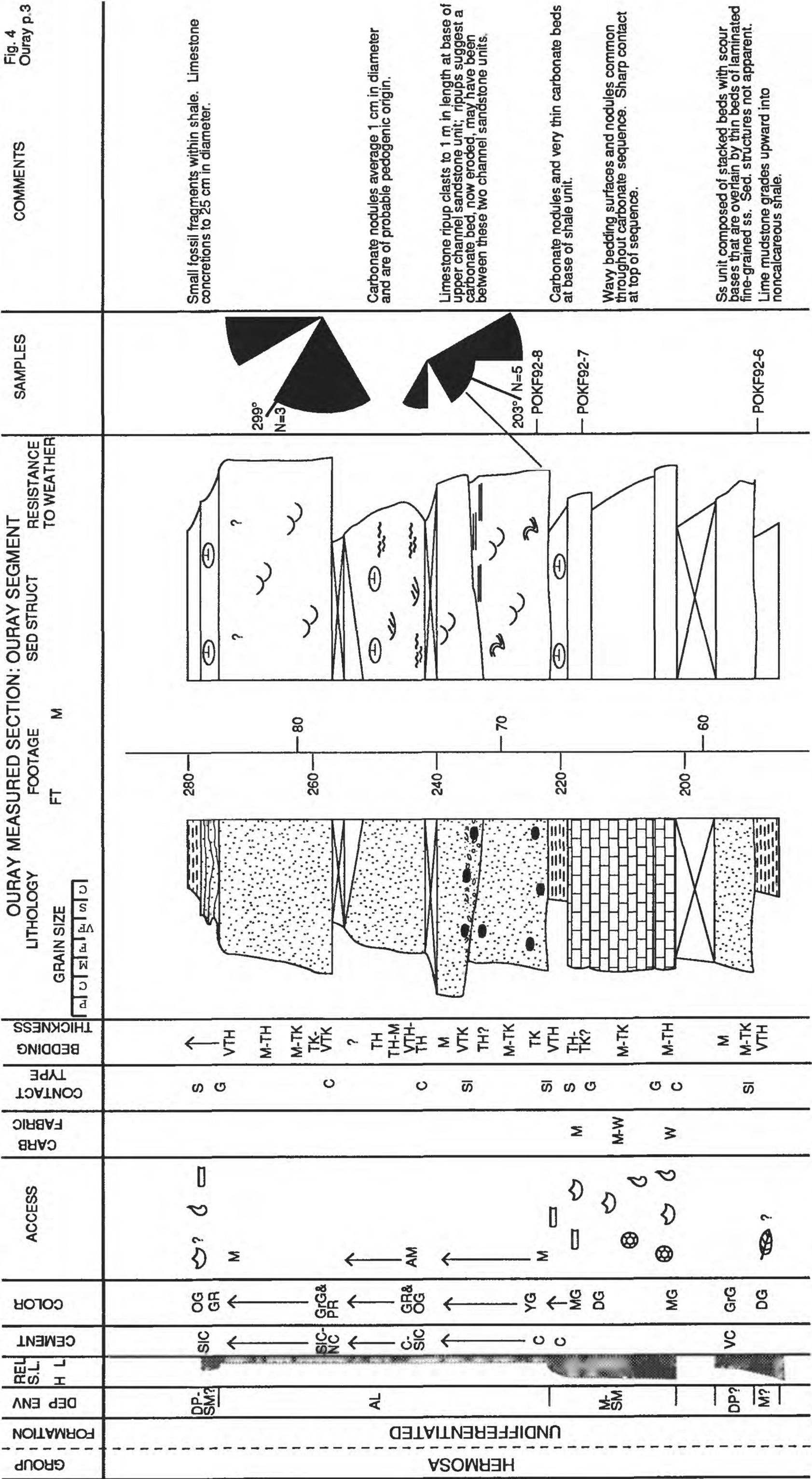
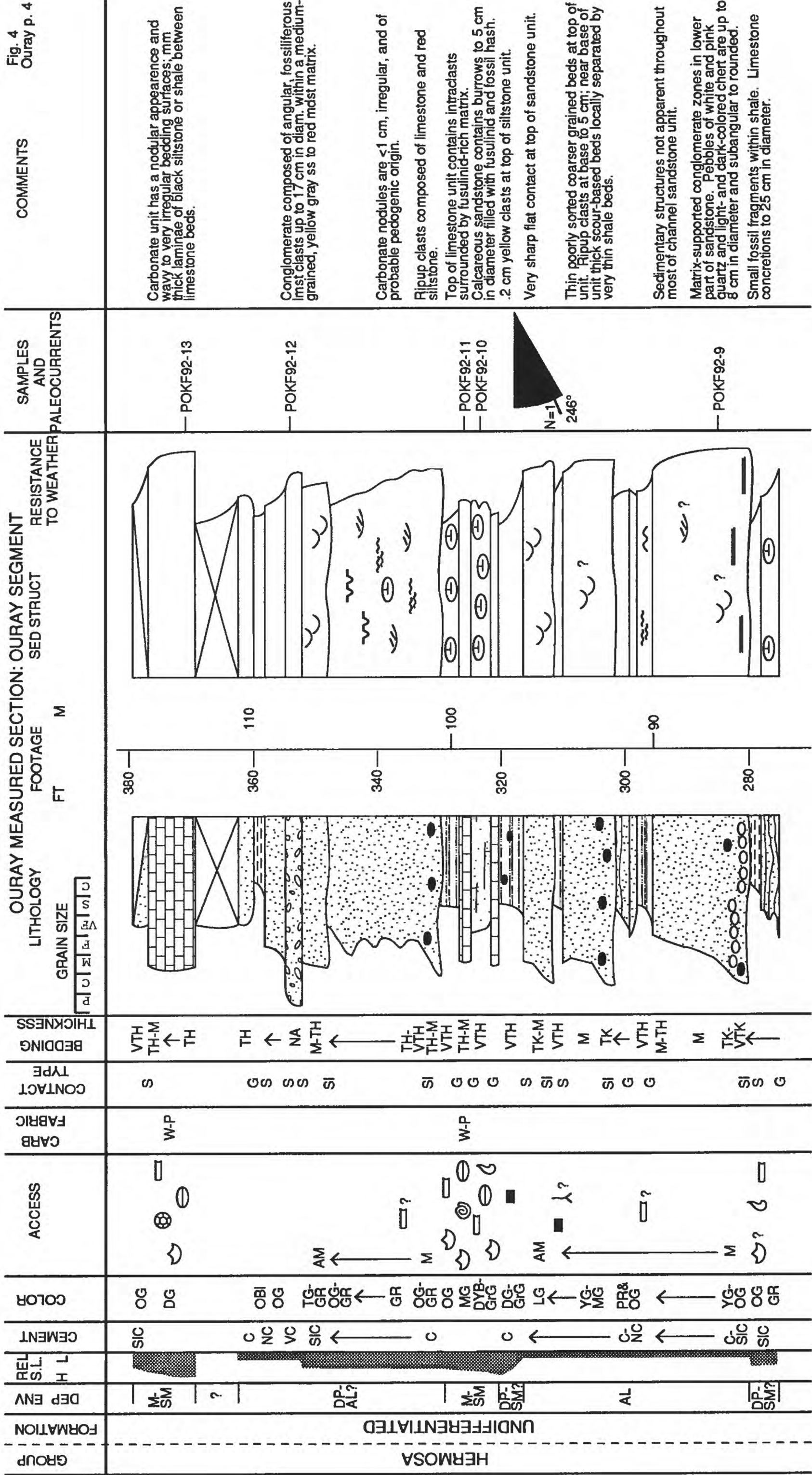




Fig. 4  
Ouray p. 4





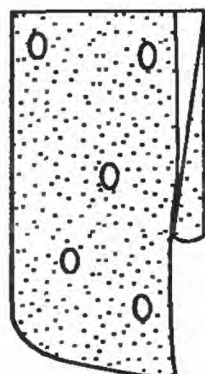

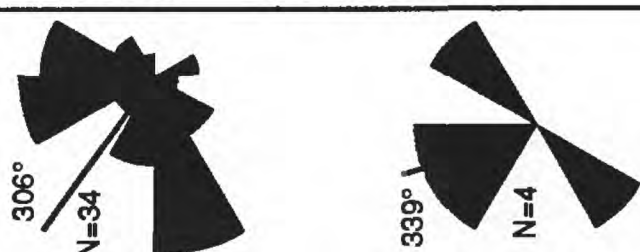
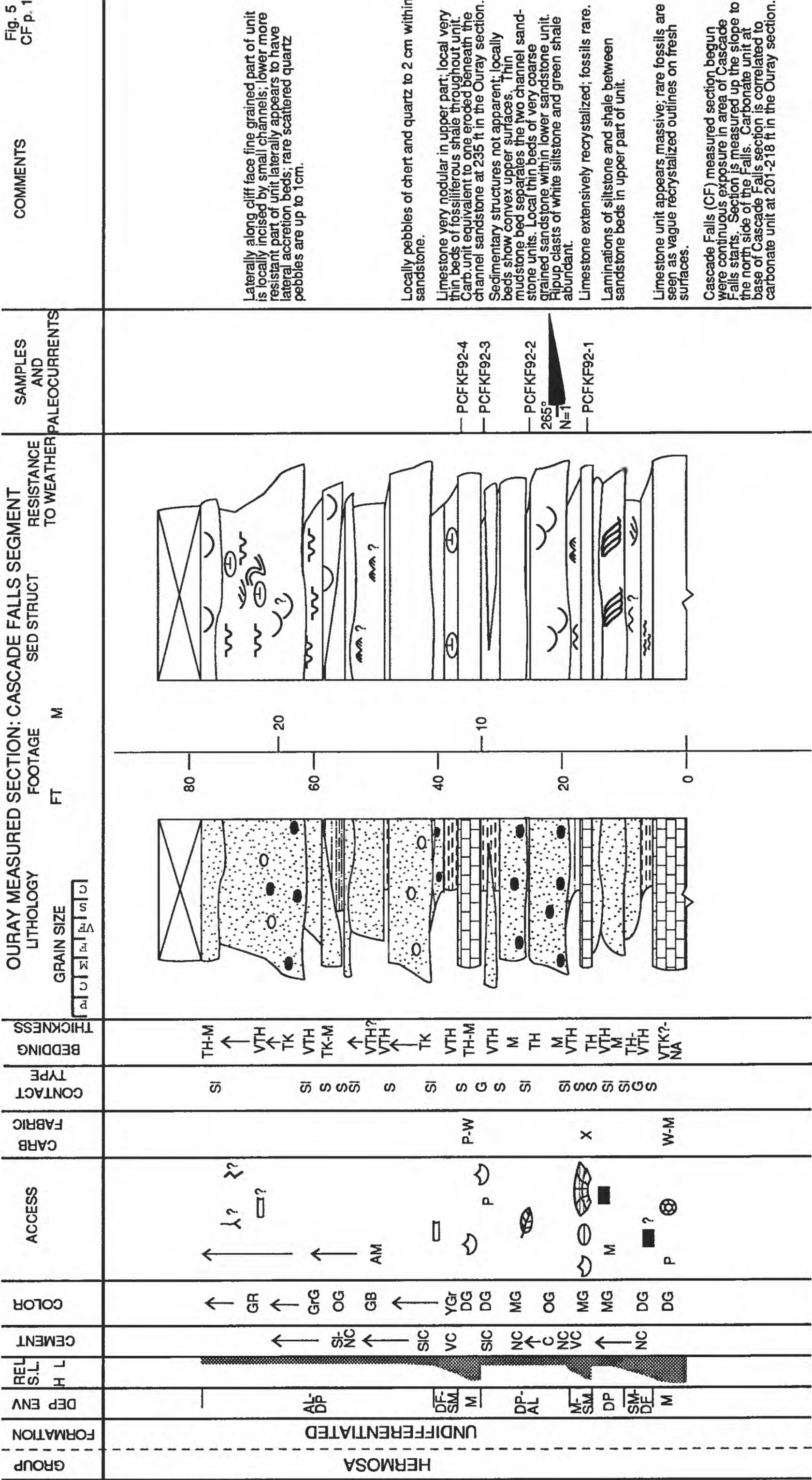
GROUP	FORMATION	DEP ENV	REL S.L.	CEMENT	COLOR	ACCESS	CARB FABRIC	CONTACT TYPE	BEDDING THICKNESS	OURAY MEASURED SECTION: OURAY SEGMENT			RESISTANCE TO WEATHER	SAMPLES AND PALEOCURRENTS	COMMENTS
										LITHOLOGY	FOOTAGE	SED STRUCT			
										GRAIN SIZE	FT	M			
										p   c   m   r   s   v					
HERMOSA	UNDIFFERENTIATED	AL	H	L	C	LG	M	S	M-TK	VTH				<p>Rose diagram to left is of all paleocurrent readings taken through the Ouray measured section and those taken through sandstone units that are up to 500 ft above the top of the Ouray section.</p> <p>Section above cliff-forming sandstone obscured by talus and vegetation. Locally exposed about 5 ft above cliff-forming sandstone is a limestone unit about 3 ft thick.</p> <p>Top of Ouray measured section</p> <p>Pebbles of chert and quartzite &lt; 1 cm are common in lower 5 ft of sandstone unit. Cliff-forming sandstone overlain by slope-forming medium- to fine-grained sandstone that is thin to very thin bedded and mottled red and green.</p>	

Fig. 5  
CF p. 1







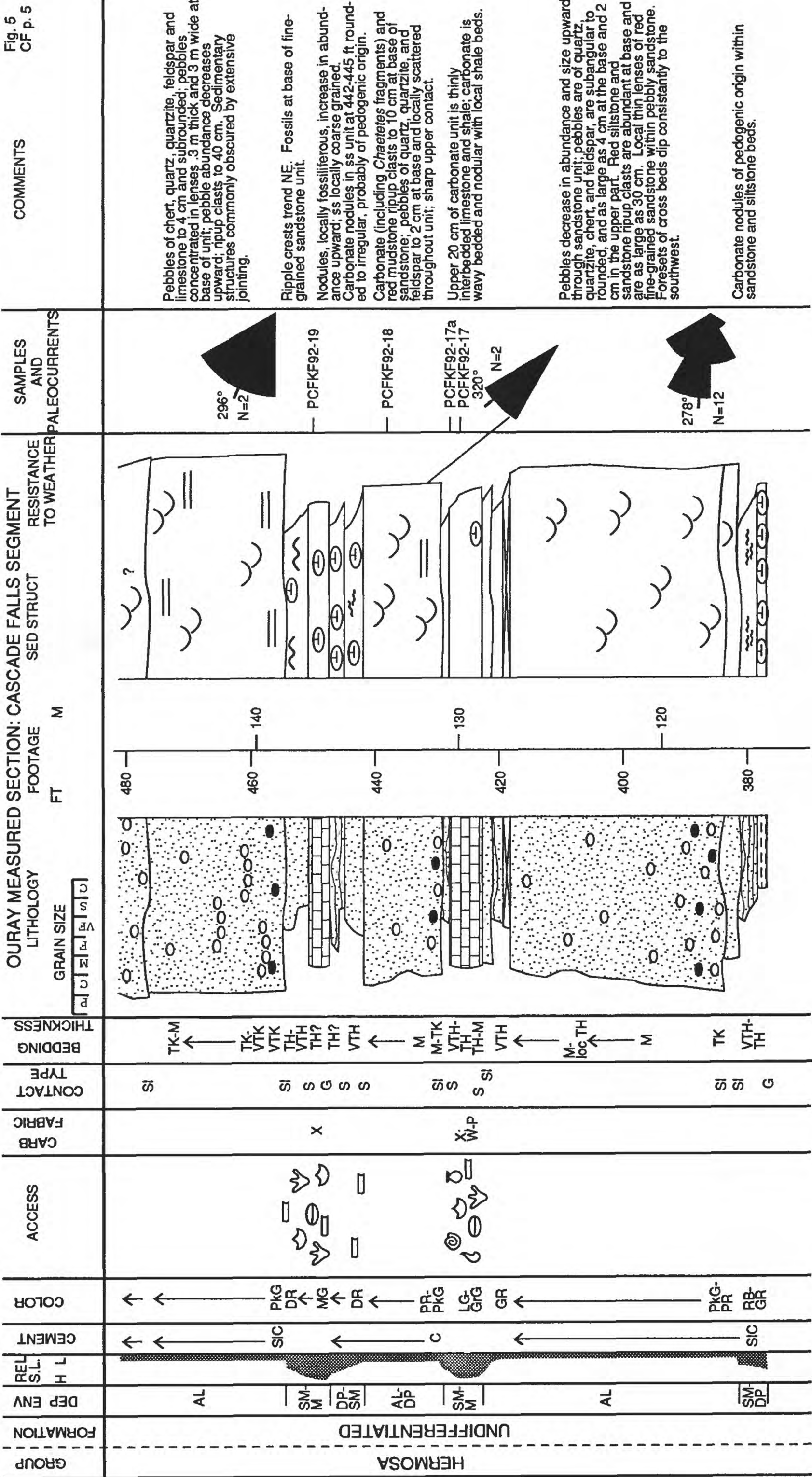
GROUP	FORMATION	DEP ENV	I T'S	CEMENT	COLOR	ACCESS	CARB FABRIC	CONTACT TYPE	BEDDING THICKNESS
1									
2									
3									
4									
5									
6									
7									
8									
9									
10									
11									
12									
13									
14									
15									
16									
17									
18									
19									
20									
21									
22									
23									
24									
25									
26									
27									
28									
29									
30									
31									
32									
33									
34									
35									
36									
37									
38									
39									
40									
41									
42									
43									
44									
45									
46									
47									
48									
49									
50									
51									
52									
53									
54									
55									
56									
57									
58									
59									
60									
61									
62									
63									
64									
65									
66									
67									

[illegible]





Fig. 5  
CF p. 5





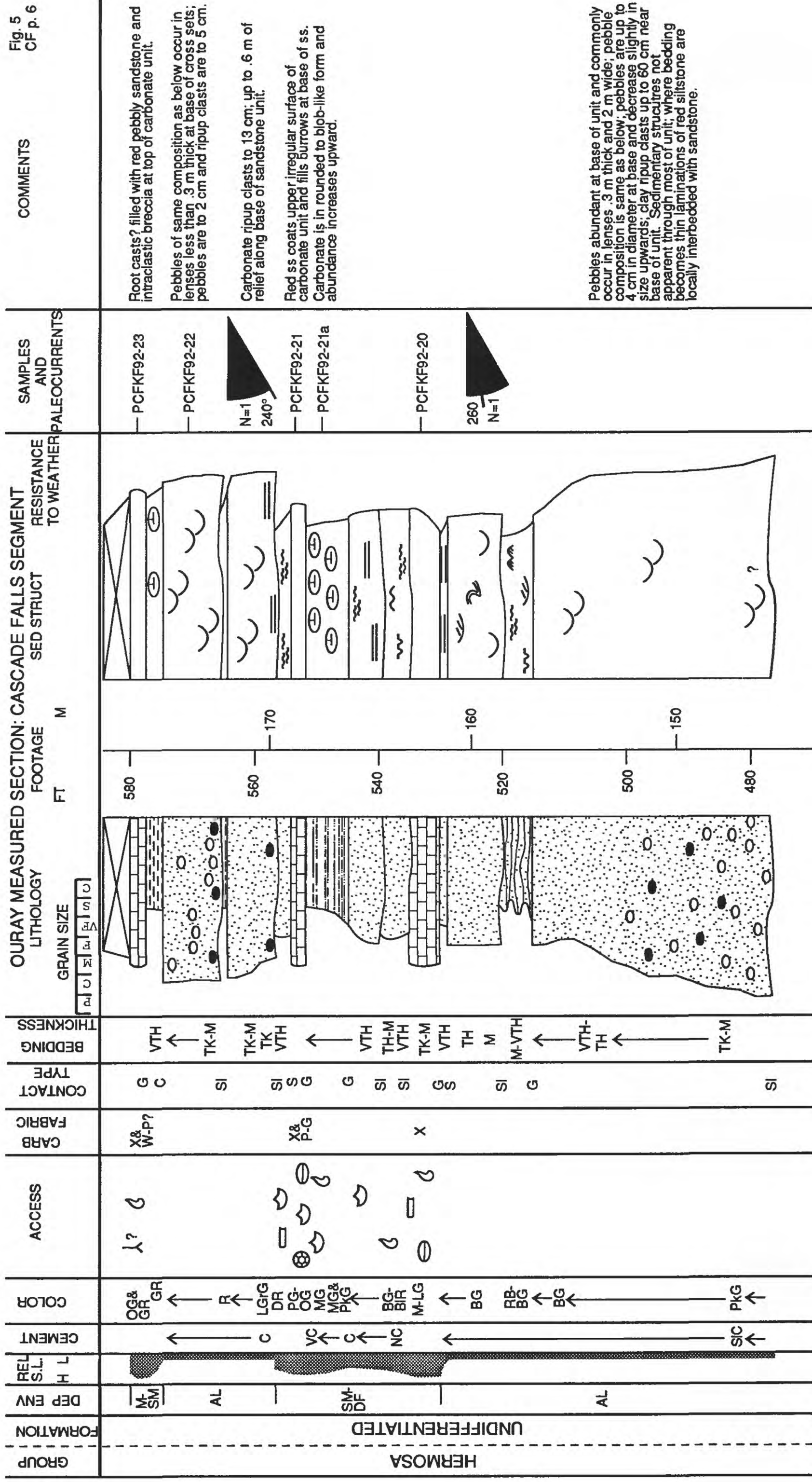






Fig. 5  
CF p. 8

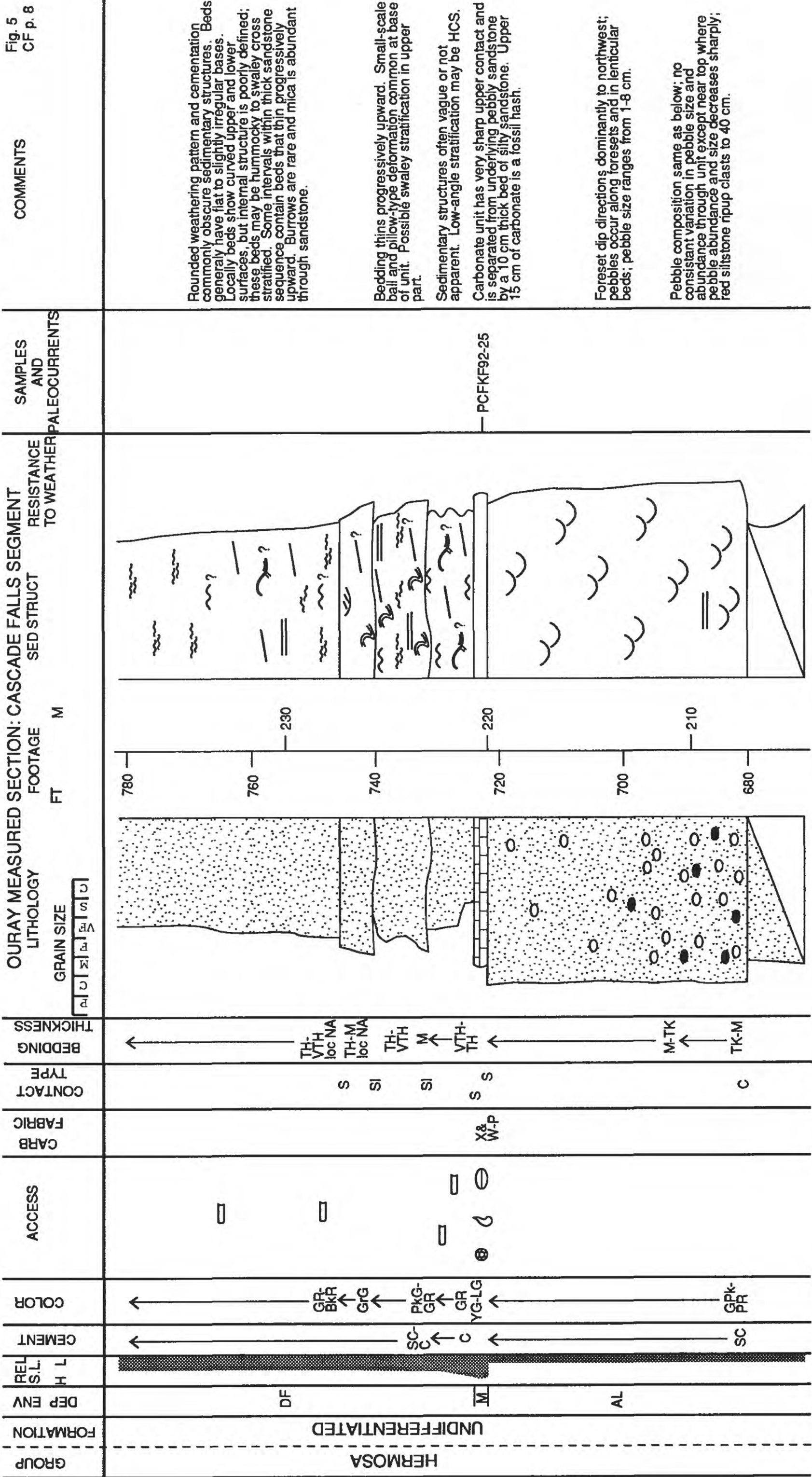


Fig. 5  
CF p. 9

GROUP	FORMATION	DEP ENV	REL S.L. H L	CEMENT	COLOR	ACCESS.	CARB FABRIC	CONTACT TYPE	BEDDING THICKNESS	OURAY MEASURED SECTION: CASCADE FALLS SEGMENT			SAMPLES AND PALEOCURRENTS	COMMENTS
										LITHOLOGY	FOOTAGE FT	SED STRUCT M		
HERMOSA	UNDIFFERENTIATED	DF	<div></div>	<div>SC-PR-GF</div> <div>C</div>	<div></div>	<div></div>		SI	<div>VTH-loc</div> <div>TH</div>	<div></div> <div>GRAIN SIZE</div> <div>p</div> <div>c</div> <div>ck</div> <div>fs</div> <div>ve</div> <div>s</div> <div>c</div>	<div></div> <div>820</div> <div>250</div> <div>800</div> <div>240</div>	<div></div>	<div></div> <div>PCFKF92-27</div> <div>PCFKF92-26</div>	<div></div> <div>Rose diagram to left shows all paleocurrent readings taken from both the Ouray and Cascade Falls sections.</div> <div>Rose diagram to left shows all paleocurrent readings from the Cascade Falls section.</div> <div>The angular unconformity between the Dolores and Hermosa rises stratigraphically to the north of the Cascade Falls line of section. About 200 ft of the Hermosa Group has been cut out by this unconformity. This 200 ft of section is only exposed on cliff faces, and it appears to consist of alternating marine and nonmarine sequences.</div> <div>An angular unconformity separates the Triassic Dolores Formation from the underlying Hermosa Group. About 300 ft to the north of the line of section, discontinuous, fossiliferous limestone beds are within the very fine grained sandstone at the top of the Hermosa. Sample PCFKF92-26 is from these limestone beds.</div> <div>Rare bioturbated to deformed beds; burrows are generally rare through this interval; Locally beds have a nodular weathering appearance and contain blebs and stringers of sparry calcite, but no apparent fossils. Sedimentary structures are the same as below.</div>





Figure 7