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**Cruise Report: *R/V ALPHA HELIX* CRUISE-173
to western Prince William Sound, Yakutat Bay, and
Glacier Bay National Park, northeastern Gulf of
Alaska, August 17 - September 3, 1993**

by ELLEN A. COWAN², ROSS D. POWELL³, PAUL R. CARLSON¹,
ROBERT E. KAYEN¹, JINKUI CAI³, KEITH C. SERAMUR²,
and SARAH D. ZELLERS⁴

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1 U.S. GEOLOGICAL SURVEY, MENLO PARK, CA
2 APPALACHIAN STATE UNIVERSITY, BOONE, NC
3 NORTHERN ILLINIOS UNIVERSITY, DE KALB, IL
4 UNIVERSITY OF TEXAS AT AUSTIN, AUSTIN, TX

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Introduction

Cruise 173 of the *R/V Alpha Helix* from August 17 until September 3, 1993 was conducted in the fjords of Prince William Sound, Yakutat Bay and Glacier Bay, in southern and southeastern Alaska. This cruise was directed by Ellen A. Cowan of Appalachian State University and Ross D. Powell of Northern Illinois University and funded by the National Science Foundation (OPP-9223990 and OPP-9223992) under the NSF project "Collaborative Research: Definition of High Resolution Seismic Facies for Interpreting Glacial Fluctuations". Cowan's portion of the project is funded through Research in Undergraduate Institutions (RUI) and active participation by three undergraduate students in the cruise and post-cruise research reflects this aspect of NSF funding. USGS participants were funded under a cooperative study agreement with the co-PI's as part of their NSF proposal. This cruise report includes a listing of data and samples collected, maps showing track lines and station locations as well as a summary of significant initial observations.

Scientific Objectives

The overall purpose of this study is to characterize seismic facies for interpreting past glacier behavior, especially during the Last Glacial Maximum (LGM). Glacier Bay is the site of the world's best documented major marine deglaciation and allows comparisons between glacier terminus stability with changing climate and marine conditions. Upper Yakutat Bay contains deposits produced by the surge of the Hubbard Glacier in 1986 and the subsequent outburst flood from Russell fjord. Historical records of glacier terminus positions provide a datum with which the history of this marine glacial system can be determined from seismic and sedimentary facies.

We are using two techniques to characterize sedimentary assemblages: high-resolution-seismic reflection surveys and coring/grab sampling of sediment. The seismic facies are interpreted from acoustic impedance contrasts imaged within the sediment column portion of high-resolution-seismic reflection profiles. Sedimentary facies, on the other hand, are determined from physical observation and analysis of bottom samples. The sediment cores and grab samples are used to interpret sedimentary facies from the seismic facies. We will construct 3-D seismic facies to establish the spatial and temporal distribution of sedimentary facies in the glacial marine environment. This facies distribution will allow us to reconstruct the behavior of the glacial systems during climatic changes

within different marine environments. Collection of sediment cores and grab samples will allow us to determine the following additional information:

a) Measurements of physical properties of glacial marine sediment using GRAPE (Gamma Ray Attenuation Porosity Evaluator) including P-wave velocity, wet bulk density and magnetic susceptibility.

b) Determination of lithofacies characteristics including the nature of contacts, color, particle size variations, and lithology. X-radiographs will help to define sedimentary structures that are not easily observed on split core surfaces.

c) Dating of sediments using Pb-210.

d) Fourier shape analysis of silt grains to interpret glacial regime and glacial conditions from glacial marine lithofacies.

e) Identify living foraminifera in glacial marine sediments and document the downcore distribution of foraminifera and associated microfossils.

We also wish to continue to improve our understanding of sedimentary processes in modern temperate glacial marine environments. CTD profiles and water sampling provide a means of accessing the influx of meltwater to the fjords and sediment dispersal processes. These measurements will also allow us to compare glacial fjord systems under different climatic regimes in coastal Alaska.

Personnel and Support

The Chief Scientist was Ellen A. Cowan from Appalachian State University and Co-Chief Scientist was Ross D. Powell from Northern Illinois University. Other scientists included Paul R. Carlson and Robert E. Kayen from the Marine Geology Office of the U.S. Geological Survey in Menlo Park, CA; Jinkui Cai, postdoctoral fellow at Northern Illinois University; Keith C. Seramur, Adjunct research associate at Appalachian State University, and Sarah D. Zellers, Ph.D. student at University of Texas at Austin. Three undergraduate students from Appalachian State University, Jenifer D. Clark, Rafael A. Gutierrez, and Christen A. Nall assisted the Scientific party. Operator of the Huntec Deep towed system was Graham B. Standen from Geoforce Consultants, Nova Scotia. David

Hogg from the U.S.G.S., Menlo Park provided technical assistance. Paul A. Jones, of the Environmental Protection Agency (EPA) joined the cruise in Glacier Bay National Park as a Marine Mammal Observer. Detailed contact information for the scientific staff is presented in Table 1.

We wish to acknowledge the Captain and crew of the *R/V Alpha Helix* who were an important part of the success of this cruise. Despite rough seas in the Gulf of Alaska, rainy weather, icebergs and currents we accomplished our scientific goals because of their hard work and dedication to the job. We especially appreciate the efforts of Captain Tom Callahan and Chief Mate Bill Rook who kept us on track and on station, no matter where it was located. We would also like to thank Tom Smith, Marine Superintendent and his staff, Institute Marine Sciences, Seward, AK for help with pre- and post-cruise logistics. We thank Carolyn Degnan of the USGS for assistance with the navigational corrections and computer plots, and the staff of the USGS Marine Facility in Redwood City, CA, for their logistical support. Monty A. Hampton is thanked for his review of this report.

Cruise Summary

Concern by the National Park Service at Glacier Bay about the affect of sound emitted by the Hunttec seismic system on marine mammals caused us to alter our cruise plan from the planned 8 days work in Glacier Bay to 4 days. We were allowed to collect Hunttec profiles within a portion of the Bay, but only during daylight hours and periods with visibility of at least 4 km. We were not permitted to run seismic lines in Johns Hopkins Inlet or at the mouth of Glacier Bay. To comply with NPS requirements we added Paul A. Jones, biologist with the U.S. EPA, to our scientific party as a marine mammal observer. While in the Bay, he looked for humpback whales and Stellar sea lions (both endangered species) and monitored behavior of other marine mammals during operation of the seismic system. His observations indicated that we encountered no humpback whales and only three Steller sea lions in the water. The ship passed one sea lion while steaming to the start of a seismic line, therefore no evasive action was necessary. The other siting of Stellar sea lions resulted in interruption of seismic operations immediately. These two animals were not seen again after searching from the ship's crow's nest for 35 min. A total of about 60 harbor seals were encountered, 11 of which were observed with no seismic gear operating at the time. The observer reported that the marine mammals did not appear to be disrupted or harassed by the Hunttec system and reacted to our presence as they did to other ships (such as cruise ships) at close range.

We modified our cruise plan to spend less time in Glacier Bay and to include 2.5 days of reconnaissance level seismic surveying and sampling in College fjord and Harriman fjord in upper Prince William Sound. A summary table of of scientific activities is presented in Table 2.

The *R/V Alpha Helix* departed from its home port, Seward Alaska on Tuesday, August 17 at 2100 Alaska Daylight Time (ADT). CTD station 1 was occupied at the mouth of Resurrection Bay at 2248 ADT for the University of Alaska.

For the duration of the cruise, the scientific party operated on two 12-hour watches from 0600 to 1800h and from 1800 to 0600h. Seismic track lines were run from approximately 0800 until 2100h during daylight and sampling stations were occupied at night.

We arrived in Port Wells at 1015 ADT, August 18 and began running seismic lines (Figure 1). At 1655 we deployed a party in an inflatable boat near Harvard Glacier terminus (College Fjord) to collect echosounder profiles. No data were collected because the transducer broke off the transom of the boat while motoring to the terminus. We occupied 6 sampling stations in College and Harriman Fjords (Figure 2) and collected approximately 140 km of seismic lines (Figure 1).

At 0335 ADT, on August 20 we departed for Cordova via Wells and Percy Passages. At 1240 ADT Paul Carlson and Dave Hogg disembarked to fly to Yakutat to prepare gear for loading. After a rough transit through the Gulf of Alaska, we arrived at the Sitka Sound Seafood Dock in Yakutat at 2035 ADT, August 21 to load gear and pick up Paul and Dave. We then steamed to the head of the fjord near the Hubbard Glacier to run seismic lines (Figure 3). Over the next 5 days we collected 391 km of seismic profiles and occupied 25 sampling stations (Figure 4). We departed for Glacier Bay at 1239 on August 26. We arrived at Bartlett Cove, Park Headquarters at 0620 ADT on August 27. The Scientific party disembarked for meetings with Mary Beth Moss, Resource Management Specialist and Jim Taggart, Marine Biologist at Park Headquarters. Paul Jones, marine mammal observer, joined the ship. We were underway for Muir Inlet by 1348 ADT on August 27. Seismic lines (Figure 5) and stations GB-1 to GB-16 (Figure 6) were collected in Muir Inlet between August 27 until August 29 at 0705 ADT when work in the West Arm of Glacier Bay began (Figure 7). Seven stations (GB-17-GB-23) were

occupied in upper West Arm (Figure 8). In Tarr Inlet, seismic profiles were collected (Figure 9) and 11 stations (GB-24-GB-34) were occupied (Figure 10). On August 21 at 1206 ADT we transited to lower Muir Inlet to collect additional seismic data and occupy 4 sampling stations (Figures 5 & 6). In total we occupied 38 stations and collected 350 km of seismic lines in Glacier Bay. We returned to Bartlett Cove on September 1 at 0700 ADT to allow 7 members of the Scientific party to disembark for the airport in Gustavus. The ship departed Bartlett Cove at 1245 ADT on September 1 and transited directly to Seward, AK, arriving at 0955 ADT on September 3, 1993.

Bathymetry and Seismic Profiling

Navigation

GPS was the primary method of navigation for track lines and stations. The ship's location, direction and speed were stored at 30 sec intervals on computer disk. GPS was found to be unreliable for short periods around 0200 ADT in Yakutat Bay and Glacier Bay. New stations were not located during these 20-30 min periods.

Bathymetry

The ship's hull-mounted 12 kHz echosounder provided bathymetric data during all seismic surveys and during sampling. These data are especially important in these glacial settings because rapid sedimentation can change the water depth by a measurable amount each year.

Seismic Profiling

The Hunttec system was used under a subcontract with Geoforce Consultants from Nova Scotia. The system operated with a deep towed sound source and internal and external hydrophones.

Operating Parameters included:

- power output 4Kv
- firing rate 0.750 sec
- sweep rate 0.250 sec

-filter setting for internal hydrophone - 0.500 - 10.0 Khz and for external hydrophone - 0.700 - 5.0 Khz

-the fish was towed beneath the surface at depths ranging from 15 to 60 m

During profiling in College fjord the transformer in the EPC recorder for the external hydrophone overheated and failed. An EPC unit from the USGS was used in its place after being picked up in Yakutat. The quality of the external record was generally good, the internal record was generally of poorer quality. Both records were affected by irregularities in the ship's 110 VAC power supply that caused regular drops to 105 -107 VAC. We anticipate that this problem will be fixed before our next cruise.

Coring and Grab Sampling

Collecting the longest cores possible was considered the primary objective of bottom sampling during this cruise. A Benthos piston corer (model 2175) with a 2.8 in (7.1 cm) diameter core barrel was used. The lengths of sediment cores obtained were shorter than we had planned. The maximum core length collected was 376 cm, but most cores were less than 250 cm. Several modifications were made to the corer including removing the poorly functioning piston. The system was then used as a gravity corer with a 750 lb weight stand. In all, 6 piston cores and 65 gravity cores were collected (Table 3). Two box cores were collected with a 1 m³ MKIII™ box corer and subsampled with core tubes. A Van Veen sampler was used to collect 22 grab samples in locations where coring was unsuccessful (Table 3).

Physical properties were measured on all cores with the GRAPE. Cores were then split, described, photographed in black and white, and videotaped in color. Core logs are in Appendix 1. The working half of each core was subsampled for the following analyses:

<u>sample type</u>	<u>sampling interval</u>
Grain size and grain shape	5 cm for uniform lithology and variable for laminated sediment
Pb-210	5 cm
Organic carbon	variable
Moisture content	50 cm
Microfossils	variable

Physical Properties Measurements using a Multi-Sensor Whole Core Sediment Logger

Cores were logged for their physical properties on a multi-sensor whole core sediment logging device, built in Great Britain by Schultheiss Geoteck, Ltd. The logger is controlled by a personal-computer driven software system developed at the U.S. Geological Survey, Branch of Pacific Marine Geology for data acquisition and instrument manipulation. The system logs sediment bulk density, compression wave sound speed, and magnetic susceptibility of whole cores. These measurements are used to develop the physical property profiles for cores presented in Appendix 2. Robert Kayen was responsible for the operation of the logging device.

Micropaleontological Sampling

Eighty-eight samples were obtained from grab samples and trigger cores to document the distribution of living infaunal and epifaunal foraminifera. For trigger cores, subsamples were taken at 1 cm intervals down to 16 cm, then every 3-5 cm. Composite samples of approximately 50 cc were taken from grab samples. Each sample was placed into a sealed plastic bag with a solution of 4% formalin, Rose Bengal, and borax to saturation as a buffer. An additional 342 samples were taken from piston and gravity cores for analysis of downcore distribution of foraminifera. These samples of between 10 and 30 cc were taken at various intervals within the cores depending on the lithology. These samples were not stained. Later at the University of Texas at Austin both sample types were soaked overnight in a solution of calgon and water. The samples were then washed over a 63 micron sieve and dried. The identification of microfossils is being carried out by Sarah D. Zellers.

CTD Casts and Water Samples

A Neil Brown Mark III B CTD profiling system with a Turner Model III fluorometer was used to obtain CTD casts at 14 Stations (Table 4). Fluorometer measurements may be used to indicate relative concentrations of particles within the water column. The salinity, temperature and fluorometer counts collected on each downcast is included in Appendix 3. On the upcast, 12 water samples were collected from various depths with a General Oceanics Rosette sampler (Table 4). Each 1.7 L Niskin bottle was subsampled and vacuum filtered on a 0.8 micron Millipore filter. Suspended sediment concentrations were calculated and plotted by depth (Appendix 4).

Preliminary Results

1. Glacial retreat records indicate that the basins in College Fjord have been deglaciated for at least 300 years, much longer than in Glacier Bay. In College Fjord, preliminary review of seismic profiles indicates much lower total sediment accumulation than in Glacier Bay fjords indicating lower sedimentation rates. This conclusion is also supported by the degree of bioturbation and by the presence of worm tubes in cores. This reconnaissance level survey has given us information on the variability within the modern temperate glacimarine environment and raises questions about the relative importance of climatological, glaciological, and bedrock lithology as controls on sedimentation.

2. Hunttec profiles from upper Yakutat Bay show a channel buried from 3 to 6 m beneath the flat fjord floor (Figure 11). The channel appears to originate from the inlet to Russell Fjord and meanders downfjord for 9 to 10 km. The channel cross section is filled with a variable amplitude chaotic seismic facies and is cut into and buried by deposits represented on the seismic profiles by medium to high amplitude, fairly continuous, parallel, horizontal reflections. We were unable to collect cores long enough to penetrate the chaotic channel fill deposits but this seismic facies could be produced by sediment gravity flows deposited by the 1986 outburst flood from Russell Fjord. We are working on bathymetric data from this cruise to compare with bathymetry collected by NOAA prior to the outburst flood to check the timing of channel formation and development. It appears that a small channel existed prior to 1986, but during the outburst event it was rapidly deepened by erosion and then completely infilled with sediment. The rapid burial of this channel aided its preservation in the glacimarine record and, therefore, may be used as an indicator of surge/outburst events in Yakutat Bay.

3. Visual descriptions and x-radiographs of cores collected from upper Yakutat Bay show an apparent regular repetition of thick couplets of diamicton and laminated/homogeneous mud. These apparent cycles occur in cores collected as far as 17 km downfjord from the Hubbard Glacier. In ice-proximal cores (such as AH93YB-GC20 in Appendix 1), the diamictons are approximately 10 cm thick and the laminated mud is greater than 30 cm thick. In cores collected from distal locations (such as AH93YB-GC12 in Appendix 1), the diamictons are from 1 to 5 cm thick and the mud is less than 20 cm thick. We are testing the hypothesis that these cycles represent annual layers formed by seasonal controls on the meltwater system. The diamicton is formed by concentration of ice-rafted debris in winter and spring when there is little suspended sediment input because the meltwater system is

shut down. Although ice rafting continues through summer, suspended sediment and gravity flows deposit thick layers of interlaminated sand and mud. If these couplets are annual they will provide an important dating tool and allow us to infer sedimentation rates at all of our coring stations to better understand glacimarine sedimentary processes. We will also be able to add significant new information on the processes of diamicton formation.

4. In Glacier Bay, Huntec profiles were collected in basins proximal to Muir and Riggs Glaciers. Cores collected from these proximal basins were composed of rhythmic sand and mud laminae and massive sand beds (Figure 12). Seismic lines and sediment cores were also collected from a distal fjord basin at the entrance to Wachusett Inlet. Cores collected from distal basins contain homogeneous mud and silt and mud laminae. These two data sets will be used to compare and contrast sedimentation in both proximal and distal glacimarine environments.

5. Seismic profiles were collected in a 1 km grid pattern over the morainal bank complex at the mouth of Muir Inlet. Both the higher resolution of the Huntec system and the close spatial distribution of these profiles will allow 3-D mapping of the several different seismic facies observed within this complex. We attempted to core the grounding-line fan facies, a stratified moraine facies and a push-moraine facies. The morainal bank complex is covered by a drape of approximately 3 to 4 m of suspension deposits. The longest core collected in this area was 270 cm and did not penetrate the sediment below the drape.

6. Approximately 96 km of high resolution seismic reflection profiles were collected in Tarr Inlet. Seismic profiles were collected along the entire length of the Inlet at a spacing of 1.0 to 1.5 km. Within 2 km of the terminus of Grand Pacific and Margerie Glaciers, the fjord floor has an irregular hummocky surface. The deposits in this proximal area are represented on the profiles by a chaotic seismic facies. The sediment fill in areas further downfjord is represented by a seismic facies that is composed of variable amplitude, fairly continuous, horizontal reflectors. A buried ridge with chaotic seismic facies at the entrance to Tarr Inlet was also observed. The longitudinal profile shows a sudden increase in water depth at the fjord mouth, indicating rapid basin filling inside the inlet.

7. Approximately 80 km of seismic reflection profiles were collected in the West Arm of Glacier Bay. Seismic profiles collected at a spacing of 1 km over the entrance to Queen Inlet show a deep fjord (400 m water depth) fan complex. The sediment source for this fan complex appears to be a major turbidity current channel system from Queen Inlet, a

hanging valley. The high resolution profiles contain evidence of small fan channels that cross the fan. Coring on this fan was unsuccessful; however, grab samples indicate that the surficial sediment on the fan consists of fine to very fine sand. Previous sampling efforts have yielded displaced shallow water benthic foraminifers, reinforcing the hypothesis of turbidity current origin for much of the fan sediment. In the area downfjord from the fan, the seismic profiles appear to indicate two different depositional sequences separated by an unconformity. The lower sequence is acoustically stratified and the upper sequence is a interlayered well stratified seismic facies and a reflection-free facies, with a layer of ponded reflection-free facies directly above the unconformity. At least three buried ridges with chaotic seismic facies, probably morainal banks, can be identified in this part of the West Arm.

8. Surface samples contain a variety of microfossils including benthic foraminifera, a few planktonic foraminifera, diatoms, mollusks (clams and gastropods), ostracodes, copepods, worm tubes, and plant material. Preliminary foraminiferal studies indicate rare, living individuals from the genera: *Cassidulina*, *Elphidium*, *Haplophragmoides*, *Nonionella*, and *Reophax*. Some of these stained individuals occur at depths down to 15 cm within certain trigger cores; however, most of the stained material occurs in shallower intervals (<8 cm) or in surface composites from the grab samplers. Stained clams and copepods are present in several samples.

Preliminary results from total assemblages (living & dead) indicate differences between Prince William Sound, Glacier Bay, and Yakutat Bay. Foraminifera and diatoms are rare to few in surface samples in Prince William Sound, but they are more common in Yakutat and Glacier Bays. Diatoms are abundant in some samples, particularly in Glacier Bay. In Yakutat Bay, foraminifera are more common in distal samples and fewer in samples closer to Hubbard and Turner Glaciers, probably due to dilution by terrigenous sediment near the glaciers.

Brief examination of samples from selected cores in Prince William Sound, Yakutat Bay, and Glacier Bay (one core from each area) indicates downcore changes in abundance and characteristics of the foraminiferal assemblages. The abundances of diatoms and other microfossils also appears to be variable. For example, in core AH93GBGC-7 from Glacier Bay, most samples are either barren, or contain very rare foraminifera and diatoms. From 175 to 200 cm, however, there are common to abundant foraminifera with rare diatoms. Below this interval, at 202-209 cm, diatoms are common, but foraminifera are rare.

9. Distinct downcore variations occur in physical properties in cores from all 3 study areas. Variations in P-wave velocity produce the strong contrasts in acoustic impedance suggesting that massive sand beds overlying mud can produce strong reflectors observed in seismic profiles. Magnetic susceptibility shows great variability in cores from Yakutat Bay and Glacier Bay but has lesser variability in Prince William Sound cores. Bulk density logs and visual core descriptions indicate that mud and diamicton layers generally have low magnetic susceptibility and sands have high magnetic susceptibility. This may prove to be another method of detecting depositional cycles in cores.

10. CTD casts from College fjord show a thick, near surface layer with relatively warm water overlying low temperature, high salinity deep water. Suspended sediment concentrations range from 3 mg/L up to nearly 300 mg/L at the most ice proximal station. Four of the five sediment concentration plots have the highest concentration in the middle of the water column between 60 and 80 m. CTD casts from Yakutat Bay show similarities to those from College Fjord although surface temperatures were generally lower and deep water slightly warmer in Yakutat Bay. Stations YB-1 through YB-4 and YB-21 (all located in upper Yakutat Bay) had a lower salinity surface layer that corresponds to the highest suspended sediment concentration in the profile. Salinity is uniform beneath the surface layer in these profiles. Station YB-7, located near the confluence of the upper Bay (Disenchantment Bay) and the more open part of Yakutat Bay has an unstable salinity profile and its highest sediment concentration occurs at 45 m depth. Concentrations at that depth of 30 mg/L indicates that suspended sediment is dispersed over long distances from the head of the fjord. CTD casts from Glacier Bay show a thin low salinity overflow with stable higher salinity water beneath. The surface overflow corresponds with peaks in near surface sediment concentration. At each of the 3 stations in Glacier Bay suspended sediment gradually increases with depth through the water column. These data provide information about suspended sediment transport in each of the 3 study areas. Our previous work in southeastern Alaska has shown that water column properties and sediment concentrations are greatly influenced by the station location with respect to tidewater glaciers and tidal stage. We are presently considering the influences of these variables on the new data.

Distribution of Samples and Data

Cores and sediment samples:

The archive half of each split core is stored in a cold room at the USGS, Marine Geology Branch at Menlo Park, CA. Subsamples from the working half of each split core were shipped to the following institutions:

Appalachian State University -- samples for grain size analysis and grain shape analysis

Northern Illinois University -- samples for Pb-210 analysis, samples for total organic carbon, and slabs collected for thin sections

USGS, Menlo Park -- samples for moisture content and unsplit cores

University of Texas, Austin -- samples for microfossil identification

Cores were x-rayed at the USGS by Jinkui Cai and Paul Carlson. Original x-ray negatives for cores from Prince William Sound, Yakutat Bay, and Muir Inlet are at ASU. X-ray negatives for the West Arm of Glacier Bay and Tarr Inlet are at NIU.

Bulk samples from box cores and Van Veen grabs were shipped to NIU. Samples collected for microfossil identification from grabs are at University of Texas.

Filters from water samples were shipped to ASU where they were weighed and suspended sediment concentrations were calculated.

Bathymetry, Hunttec Profiles and Navigation

All echosounder profiles and Hunttec profiles were microfilmed at the USGS and copies were sent to NIU and ASU. The original 12 kHz record is stored at the USGS and the Hunttec record is at NIU. All ship's navigation data were taken to the USGS where the track lines were computer plotted.

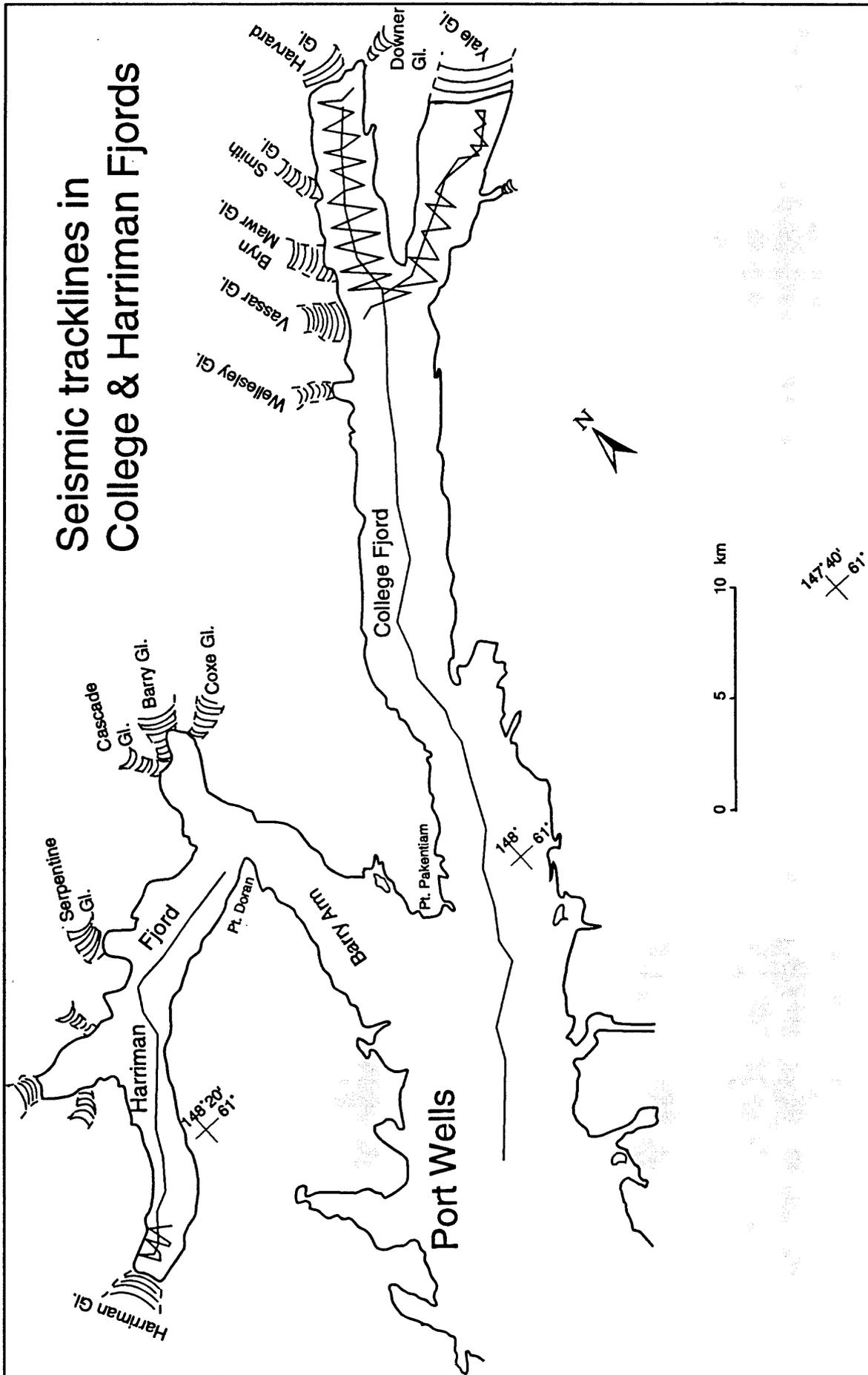


Figure 1. Seismic tracklines in College and Harriman Fjords

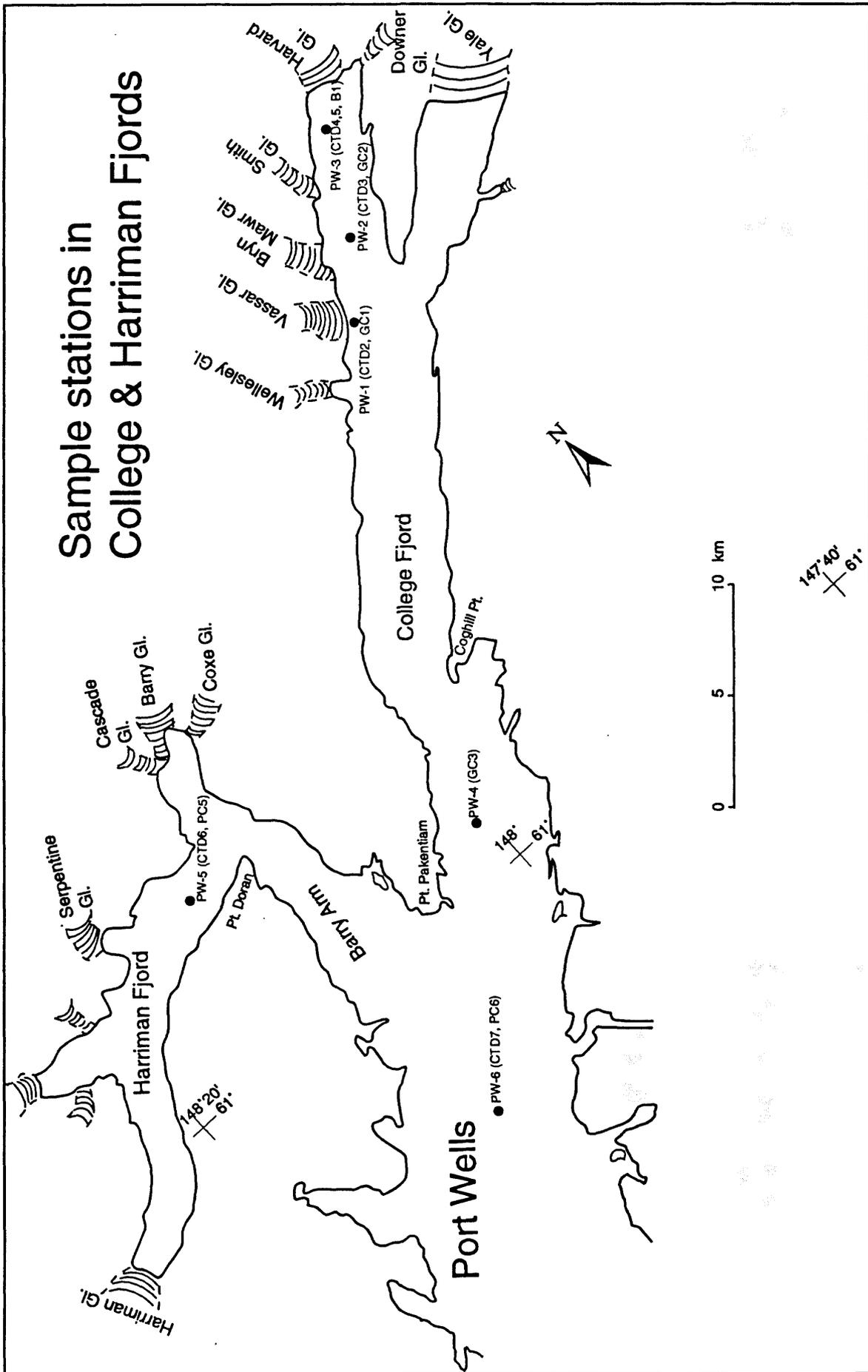


Figure 2. Sample stations in College and Harriman Fjords

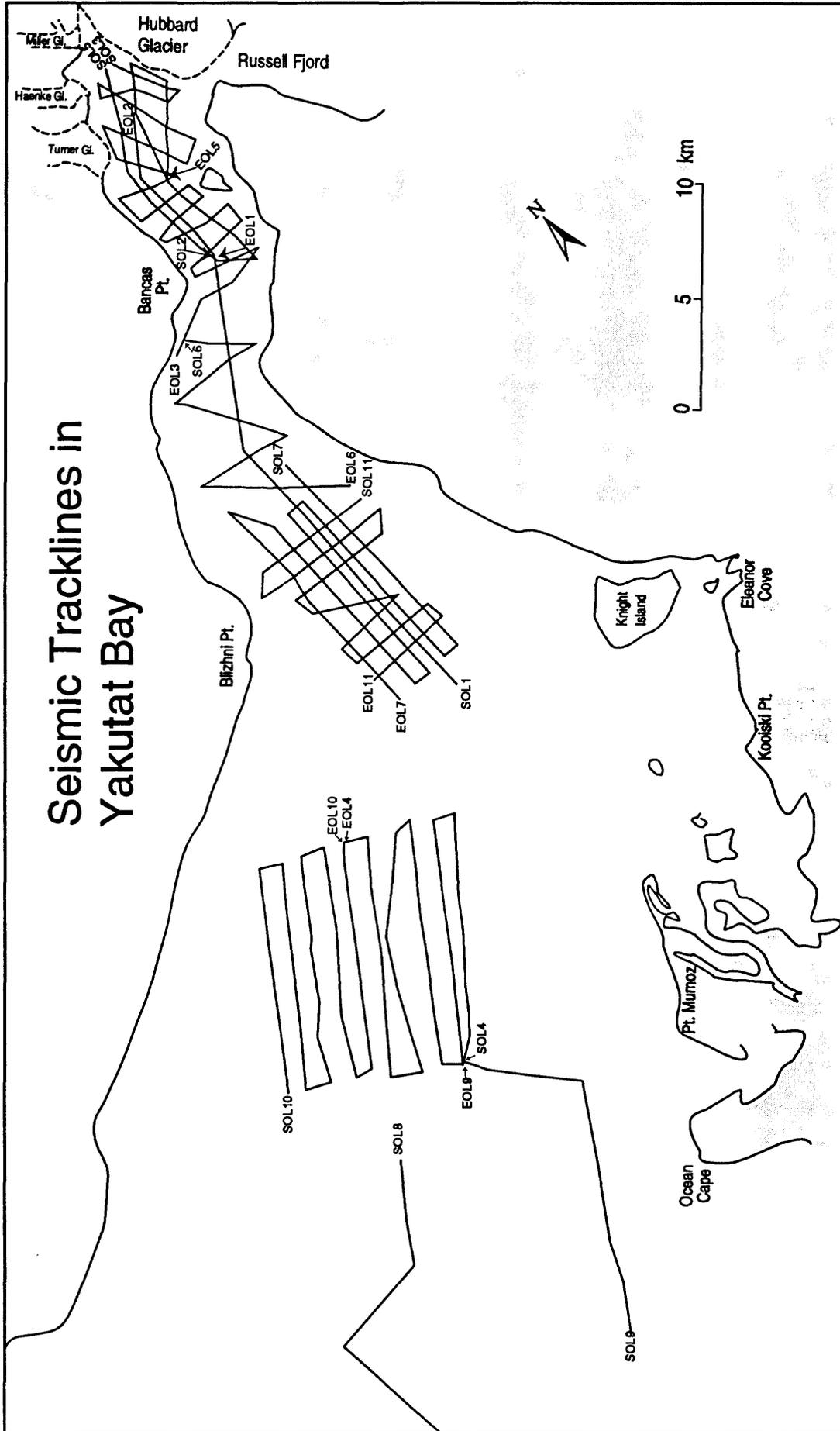


Figure 3. Seismic tracklines in Yakutat Bay

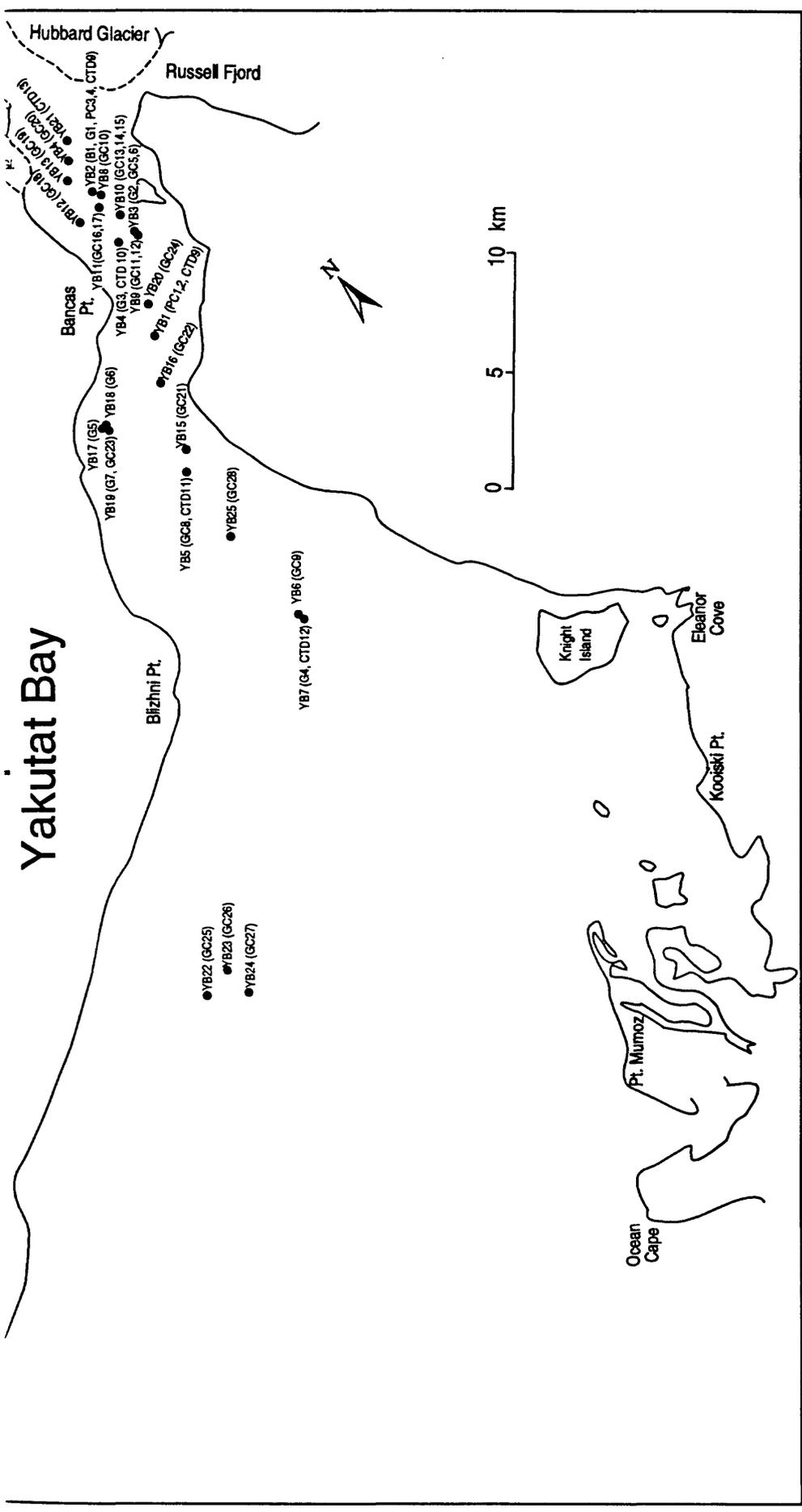


Figure 4. Sample stations in Yakutat Bay

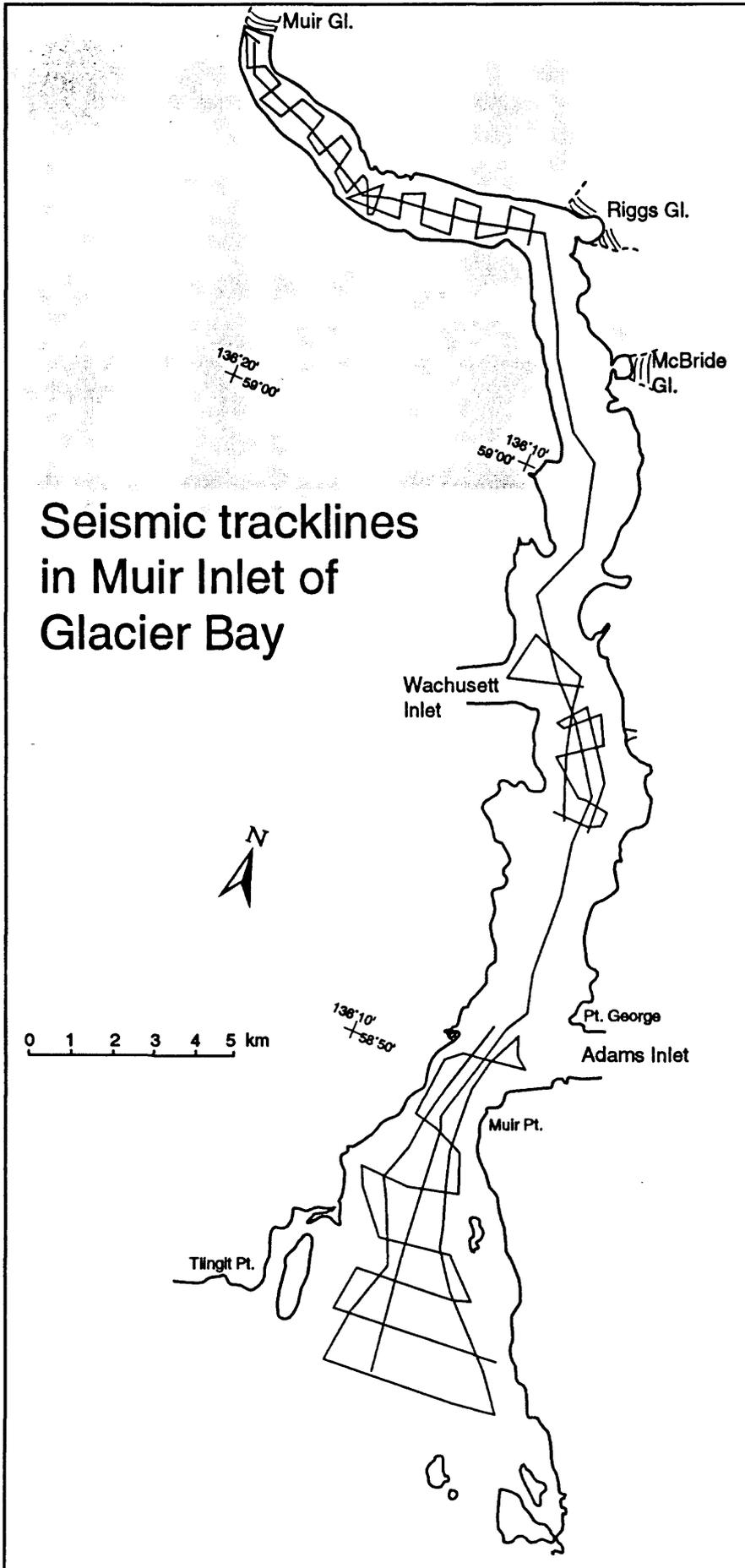


Figure 5. Seismic tracklines in Muir Inlet of Glacier Bay

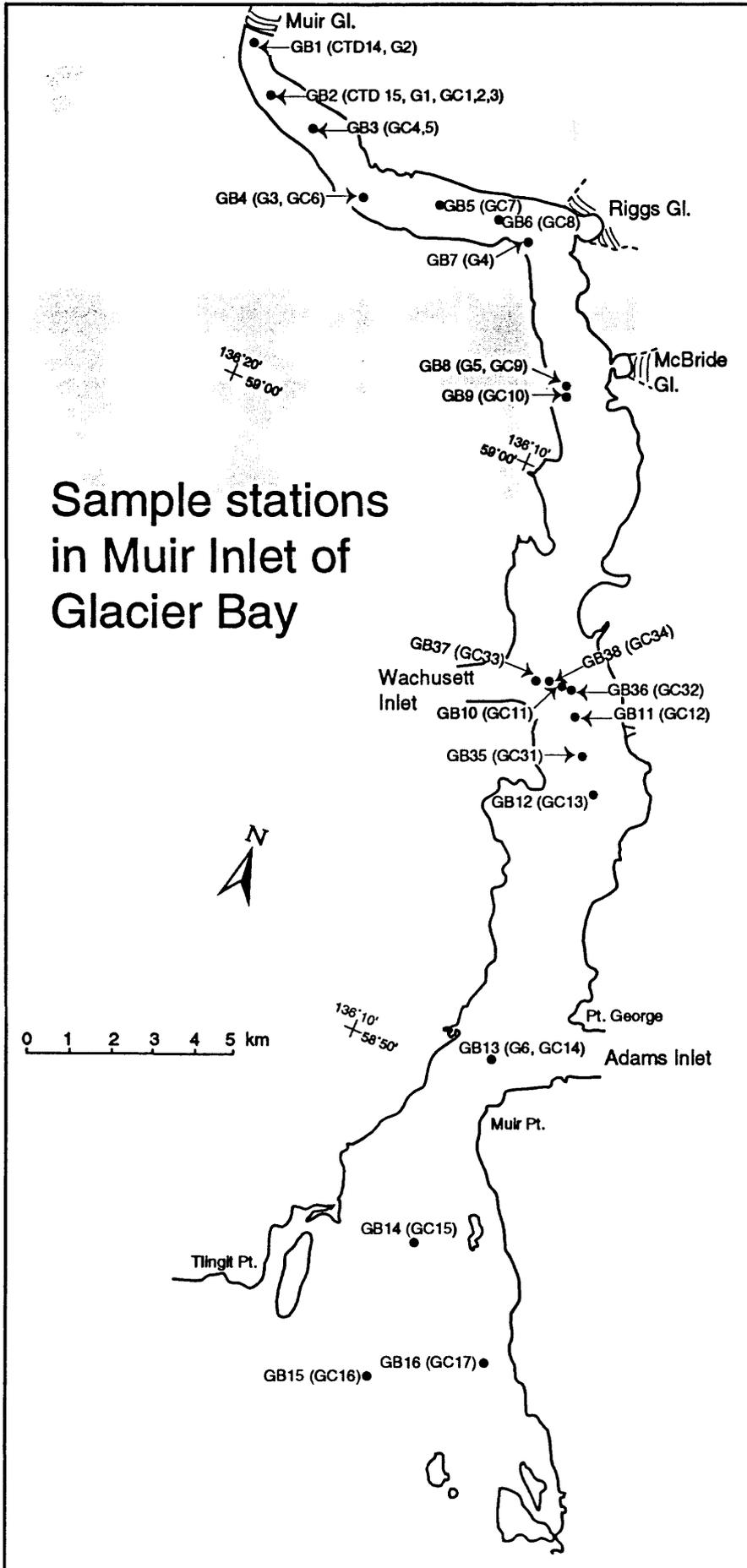


Figure 6. Sample stations in Muir Inlet of Glacier Bay

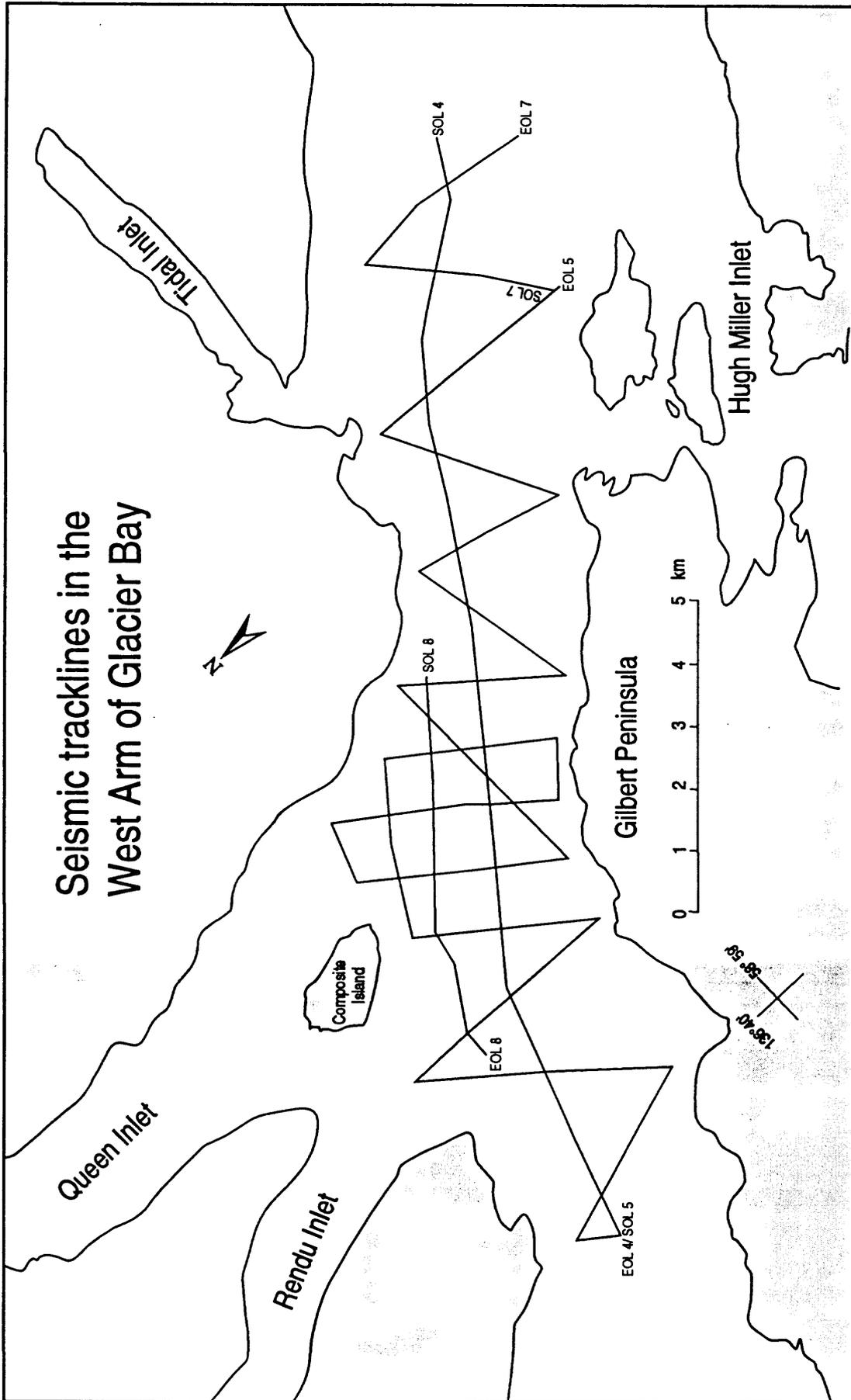


Figure 7. Seismic tracklines in the West Arm of Glacier Bay

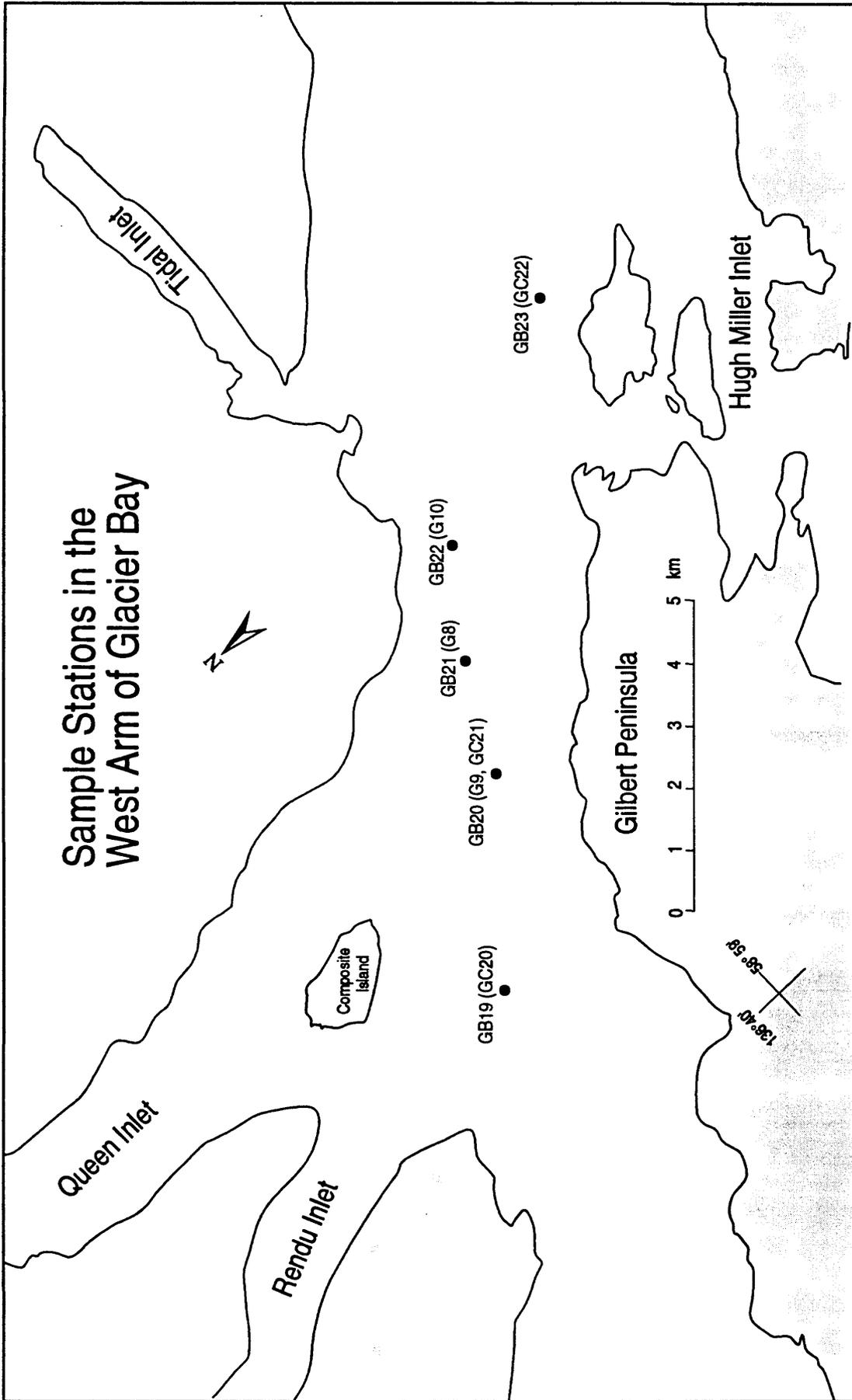


Figure 8. Sample stations in the West Arm of Glacier Bay

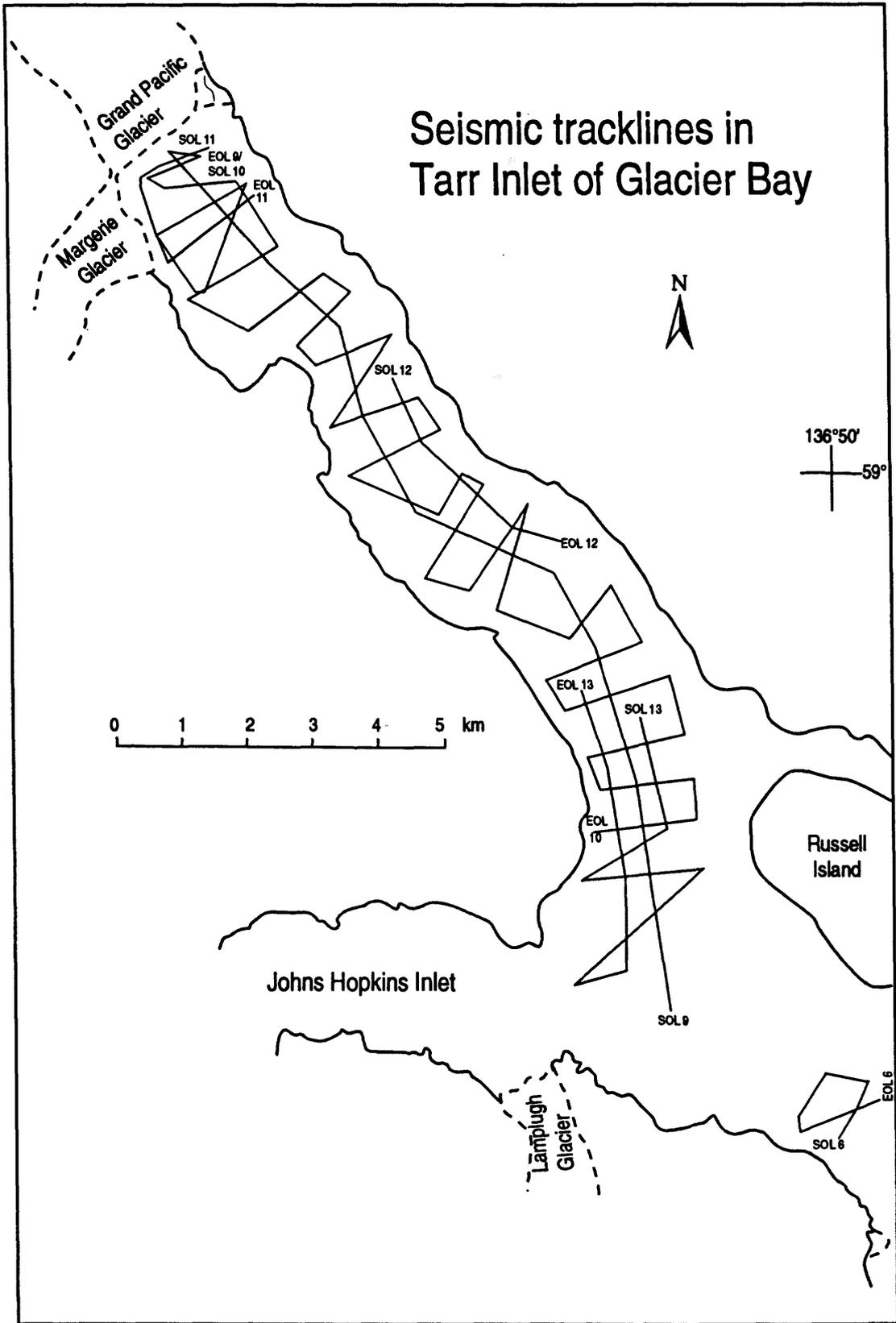


Figure 9. Seismic tracklines in Tarr Inlet of Glacier Bay

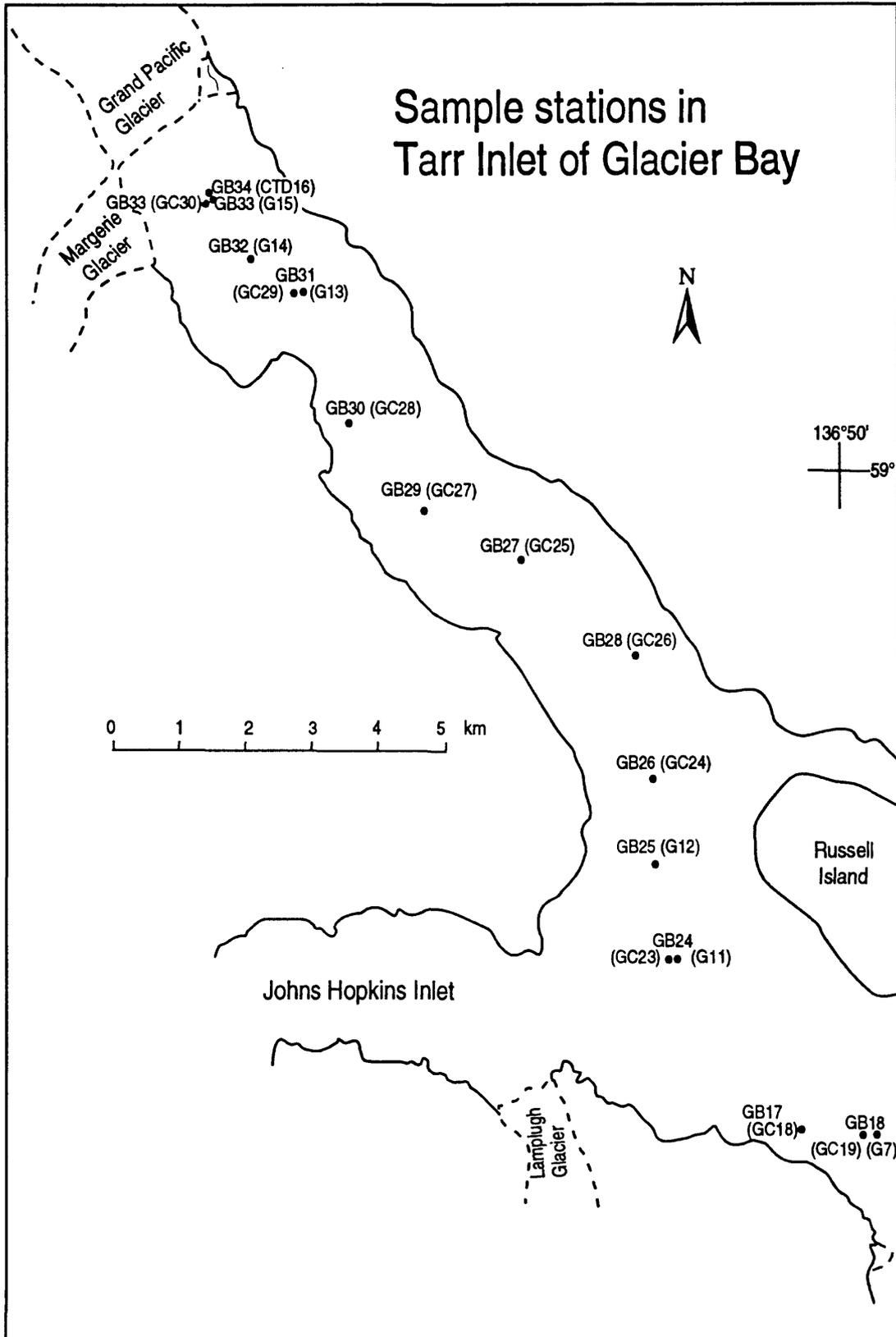


Figure 10. Sample stations in Tarr Inlet of Glacier Bay

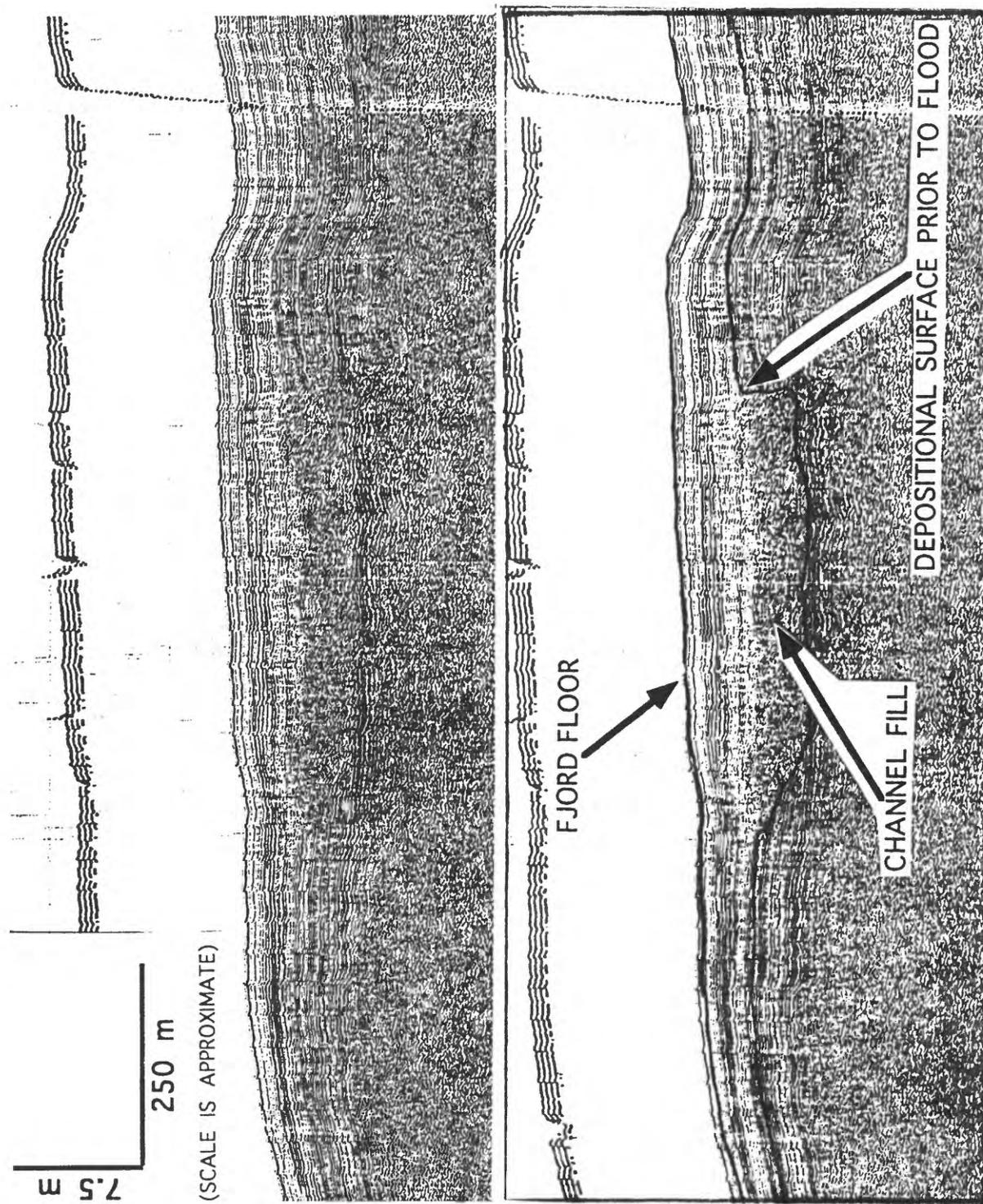


Figure 11. Seismic profile from Yakutat Bay showing buried channel.



Figure 12. Photograph of core from Muir Inlet showing interlaminated sand and mud. This 151 cm long gravity core was collected from 224 m water depth in Muir Inlet. See Figure 6 for location of core GC3 in Muir Inlet.

Table 1 - The Scientific Party

Ellen A. Cowan, Chief Scientist Department of Geology Appalachian State University Boone, North Carolina 28068	Phone: 704-262-2260 FAX: 704-262-2127
Ross D. Powell, Co-Chief Scientist Department of Geology Northern Illinois University DeKalb, Illinois 60115	Phone: 815-753-0523 FAX: 815-753-1945
Paul R. Carlson, Research Geologist Office of Marine Geology U.S. Geological Survey 345 Middlefield Road, MS 999 Menlo Park, California 94025	Phone: 415-354-3066 FAX: 415-354-3191
Robert E. Kayen, Research Civil Engineer Office of Marine Geology U.S. Geological Survey 345 Middlefield Road, MS 999 Menlo Park, California 94025	Phone: 415-354-3036 FAX: 415-354-3191
Keith C. Seramur, Geologist Department of Geology Appalachian State University Boone, North Carolina 28608	Phone: 704-328-2991 FAX: 704-322-2268
Jinkui Cai, Geologist Department of Geology Northern Illinois University Dekalb, Illinois 60115	Phone: 815-753-6272 FAX: 815-753-1945
Sarah D. Zellers, Geologist Department of Geological Sciences University of Texas at Austin Austin, Texas 78712	Phone: 512-471-6955 FAX: 512-471-9425
Graham B. Standen, Geoforce Consultants Bedford Institute of Oceanography P.O. Box 696 Dartmouth, Nova Scotia B2Y 3Y9	Phone: 902-463-0932 FAX: 902-464-9602
Undergraduate Students from Appalachian State University: Jenifer D. Clark Rafael A. Gutierrez Christen A. Nall	
David Hogg, Marine Technician U.S. Geological Survey	
Paul A. Jones, Marine Mammal Observer U.S. EPA, San Francisco, California	

Table 2. Time-table of scientific activities.

Date (1993)	Time (ADT)	Activity
August, 14-16		Mobilize shipboard equipment onboard the <i>R/V Alpha Helix</i> , Seward, AK
August 17	2100	Depart Seward, AK
August 17	2248	CTD cast, mouth of Resurrection Bay
August 18	1015	Arrive at Port Wells, AK. Begin survey of College Fjord.
August 18	1655	Echosounder Profiles of Harvard Glacier Terminous, survey College Fjord.
August 20	0336	Depart College Fjord for Cordova, AK
August 20	1240	Arrive at Cordova, AK, Carlson and Hogg fly to Yakutat, AK. Depart Cordova for Yakutat Bay.
August 21	2035	Arrive at Yakutat, AK, survey Yakutat Bay.
August 26	1239	Depart Yakutat Bay for Glacier Bay Nat'l Park.
August 27	0620	Arrive at Bartlett Cove, AK.
August 27	1348	Begin survey of Muir inlet, West Arm, and Tarr Inlet, Glacier Bay Nat'l. Park
September 1	0700	Return to Bartlett Cove, AK
September 1	1245	Depart Bartlett Cove for transit to Seward, AK.
September 3	0955	Arrive at Seward, AK
September 3-4		Demobilize operations on <i>R/V Alpha Helix</i>

Table 3. Core and Grab Samples

STATION NUMBER	SAMPLE NUMBER	DEPTH (meters)	CORE LENGTH (cm)	LATITUDE	LONGITUDE
College Fjord					
PW-1	AH93PWGC-1	230	169	61°11.28	147°48.09
PW-2	AH93PWGC-2	217	209	61°13.25	147°46.38
PW-3	AH93PWB-1	185	63	61°15.25	147°44.11
PW-4	AH93PWGC-3	104	48	61°01.07	148°00.41
PW-5	AH93PWPC-5	124	221	61°04.13	148°13.31
PW-6	AH93PWPC-6	390	234	60°55.46	148°08.46
Yakutat Bay					
YB-1	*AH93YBPC-1	246	90	59°55.24	139°37.11
	AH93YBPC-2	240	121	59°55.24	139°36.54
YB-2	AH93YBPC-3	235	0	59°58.87	139°33.69
	AH93YBG-1	233		59°58.83	139°33.95
	AH93YBB-1	233		59°58.80	139°34.25
YB-3	AH93YBPC-4	233	137	59°58.93	139°33.83
	AH93YBGC-5	237	98	59°57.50	139°33.86
	AH93YBGC-6	236	176	59°57.59	139°33.85
	AH93YBG-2	236		59°57.54	139°33.86
YB-4	AH93YBGC-7	234	277	59°57.68	139°34.76
	AH93YBG-3	234		59°57.68	139°34.83
YB-5	AH93YBGC-8	254	288	59°53.11	139°40.35
YB-6	AH93YBGC-9	86	40	59°48.30	139°41.51
YB-7	AH93YBG-4	82		59°48.46	139°41.51
YB-8	AH93YBGC-10	232	202	59°58.76	139°33.99
YB-9	*AH93YBGC-11	236	123	59°57.54	139°33.80

Table 3. Core and Grab Samples

STATION NUMBER	SAMPLE NUMBER	DEPTH (meters)	CORE LENGTH (cm)	LATITUDE	LONGITUDE
	AH93YBGC-12	236	152	59°57.54	139°33.85
YB-10	*AH93YBGC-13	234	94	59°58.05	139°33.70
	*AH93YBGC-14	234	53	59°58.06	139°33.70
	AH93YBGC-15	234	201	59°58.08	139°33.74
YB-11	AH93YBGC-16	232	80	59°58.59	139°34.19
	*AH93YBGC-17	232	60	59°58.58	139°34.16
YB-12	AH93YBGC-18	229	230	59°58.70	139°35.34
YB-13	AH93YBGC-19	227	236	59°59.54	139°34.31
YB-14	AH93YBGC-20	225	224		
YB-15	AH93YBGC-21	252	255	59°53.20	139°39.83
YB-16	AH93YBGC-22	248	282	59°54.69	139°38.36
YB-17	AH93YBG-5	61		59°54.86	139°41.74
YB-18	AH93YBG-6	90		59°54.80	139°41.34
YB-19	AH93YBG-7	95		59°54.78	139°41.18
	AH93YBGC-23	100	197	59°54.79	139°41.17
YB-20	AH93YBGC-24		150	59°56.17	139°35.99
YB-22	AH93YBGC-25	75	277	59°44.00	139°58.11
YB-23	AH93YBGC-26	80	280	59°44.09	139°56.52
YB-24	AH93YBGC-27	101	278	59°43.30	139°55.62
YB-25	AH93YBGC-28	217	122	59°51.03	139°41.47
Glacier Bay					
GB-1	*AH93GBG-2	190		59°05.11	136°22.45
GB-2	AH93GBGC-1	225		59°04.65	136°21.58
	AH93GBG-1	225		59°04.69	136°21.48
	*AH93GBGC-2	225	135	59°04.68	136°21.49
	AH93GBGC-3	224	150	59°04.69	136°21.55

Table 3. Core and Grab Samples

STATION NUMBER	SAMPLE NUMBER	DEPTH (meters)	CORE LENGTH (cm)	LATITUDE	LONGITUDE
GB-3	*AH93GBGC-4	238	50	59°04.37	136°19.84
	AH93GBGC-5	238	98	59°04.35	136°19.84
GB-4	AH93GBG-3	230		59°03.59	136°17.54
	AH93GBGC-6	228	244	59°03.58	136°17.57
GB-5	AH93GBGC-7	270	240	59°03.68	136°15.24
GB-6	AH93GBGC-8	273	118	59°03.73	136°13.28
GB-7	AH93GBG-4	261		59°03.79	136°12.17
GB-8	AH93GBG-5	240		59°01.53	136°09.64
	AH93GBGC-9	245	144	59°01.44	136°09.66
GB-9	AH93GBGC-10	250	97	59°01.23	136°09.54
GB-10	AH93GBGC-11	290	136	58°56.57	136°06.89
GB-11	AH93GBGC-12	305	245	58°56.16	136°06.22
GB-12	AH93GBGC-13	310	327	58°54.96	136°04.98
GB-13	AH93GBG-6	181		58°50.21	136°05.58
	AH93GBGC-14	180	220	58°50.17	136°05.53
GB-14	AH93GBGC-15	70	254	58°46.89	136°06.19
GB-15	AH93GBGC-16	137	270	58°44.50	136°06.21
GB-16	AH93GBGC-17	186	376	58°45.30	136°02.78
GB-17	AH93GBGC-18	218	268	58°53.32	136°50.70
GB-18	AH93GBGC-19	280	25	58°53.26	136°44.53
	AH93GBG-7	308		58°53.27	136°49.28
GB-19	AH93GBGC-20	426		58°52.08	136°36.09
GB-20	AH93GBGC-21	424	0	58°50.80	136°33.10
	AH93GBGC-21	425	0	58°50.63	136°32.79
	AH93GBG-9	424		58°50.78	136°33.08

Table 3. Core and Grab Samples

STATION NUMBER	SAMPLE NUMBER	DEPTH (meters)	CORE LENGTH (cm)	LATITUDE	LONGITUDE
GB-21	AH93GBG-8	434		58°49.91	136°30.50
GB-22	AH93GBGC-9	434	0	58°49.29	136°28.84
	AH93GBG-10	434		58°49.24	136°28.58
GB-23	AH93GBGC-22	276	145	58°46.80	136°26.10
GB-24	AH93GBGC-23	376	0	58°55.09	136°53.33
	AH93GBGC-23	377	0	58°55.06	136°53.20
	AH93GBG-11	378		58°55.08	136°53.15
	AH93GBGC-23	378	0	58°55.04	136°53.15
GB-25	AH93GBG-12	375		58°56.02	136°53.59
GB-26	AH93GBGC-24	340	177	58°56.76	136°53.58
GB-27	AH93GBGC-25	339	40	58°59.09	136°56.22
GB-28	AH93GBGC-26	245	55	58°58.12	136°53.96
GB-29	AH93GBGC-27	331	171	58°59.64	136°58.12
GB-30	AH93GBGC-28	318	282	59°00.57	136°59.55
GB-31	AH93GBGC-29	296		59°01.79	137°00.67
	AH93GBG-13	296		59°01.77	137°00.55
GB-32	AH93GBG-14	284		59°02.09	137°01.46
GB-33	AH93GBG-15	258		59°02.72	137°02.32
	AH93GBGC-30	265	0	59°02.69	137°02.34
GB-35	AH93GBGC-31	305	232	58°55.53	136°05.74
GB-36	AH93GBGC-32	291	152	58°56.58	136°06.76
GB-37	AH93GBGC-33	150	114	58°56.47	136°07.70
GB-38	AH93GBGC-34	287	56	58°56.55	136°07.24

* = Cores not split, described, or archived

GC = gravity core

B = box core

PC = piston core

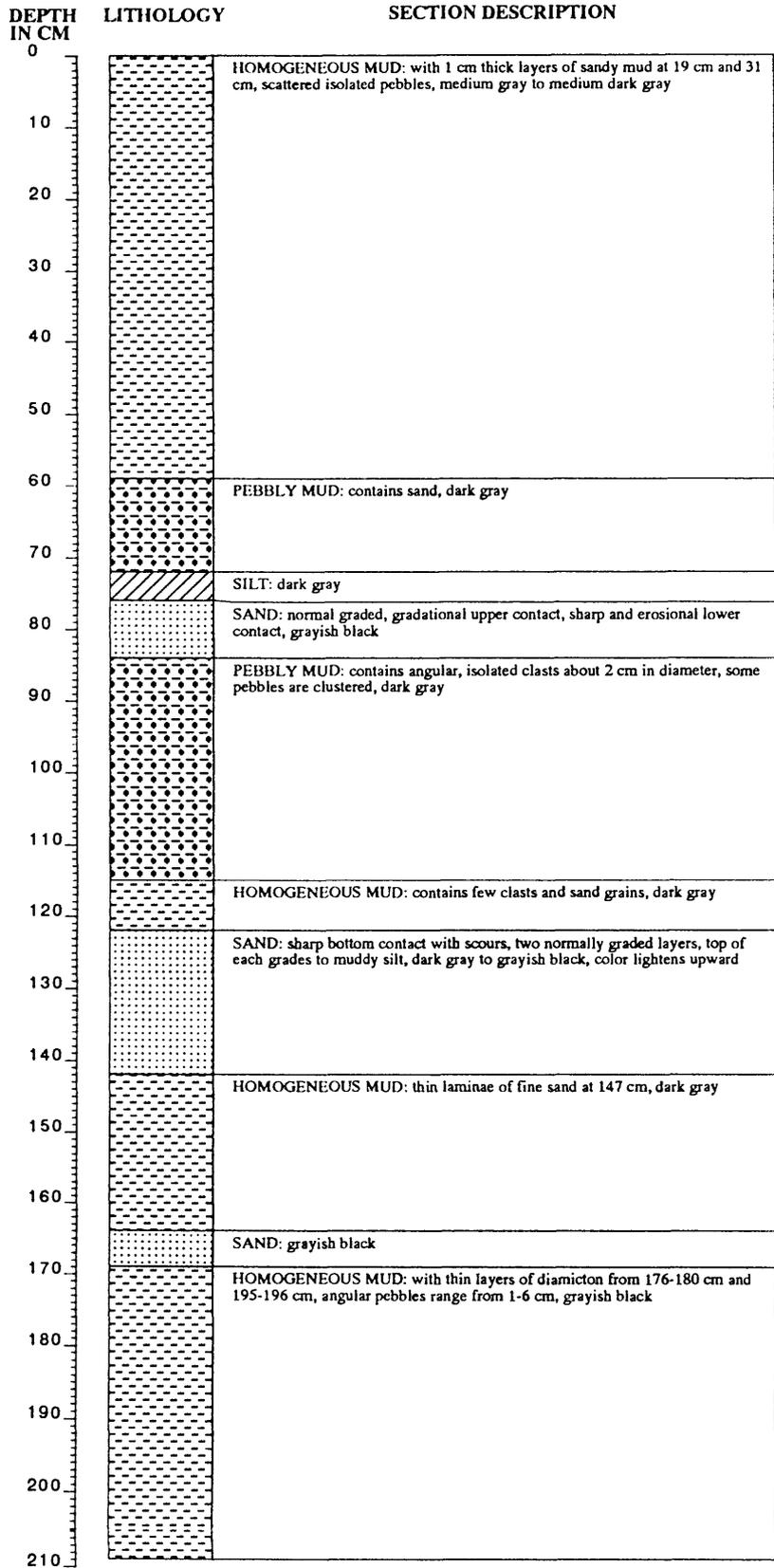
G = van veen grab

Table 4. CTD Casts and Water Samples

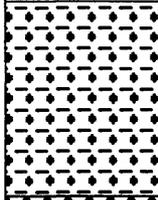
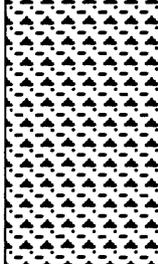
LOCATION/ STATION #	CAST	DEPTH (meters)	WATER SAMPLES (depth- meters)
College Fjord			
PW-1	2	190	5,10,20,30,40,50,60 100,145,160,190
PW-2	3	211	0,10,20,35,50,70, 110,130,150,190,202
PW-3	4	180	
PW-3	5	180	0,6.5,10,35,50,60,75, 85,110,125,150,182
PW-5	6	118	0,8,10,20,30,40,50,70, 80,92,100,118
Port Wells			
PW-6	7	365	0,10,20,40,55,100, 150,175,200,250,350, 361
Yakutat Bay			
YB-1	8	230	0,25,50,75,100,125, 150,160,180,190,200 230
YB-2	9	215	0,15,20,25,35,45,50, 60,75,100,146
YB-4	10	219	0,8,10,20,50,80,110, 130,150,170,200,219
YB-5	11	235	3.2,10,20,30,50,60, 75,100,125,150, 200,233

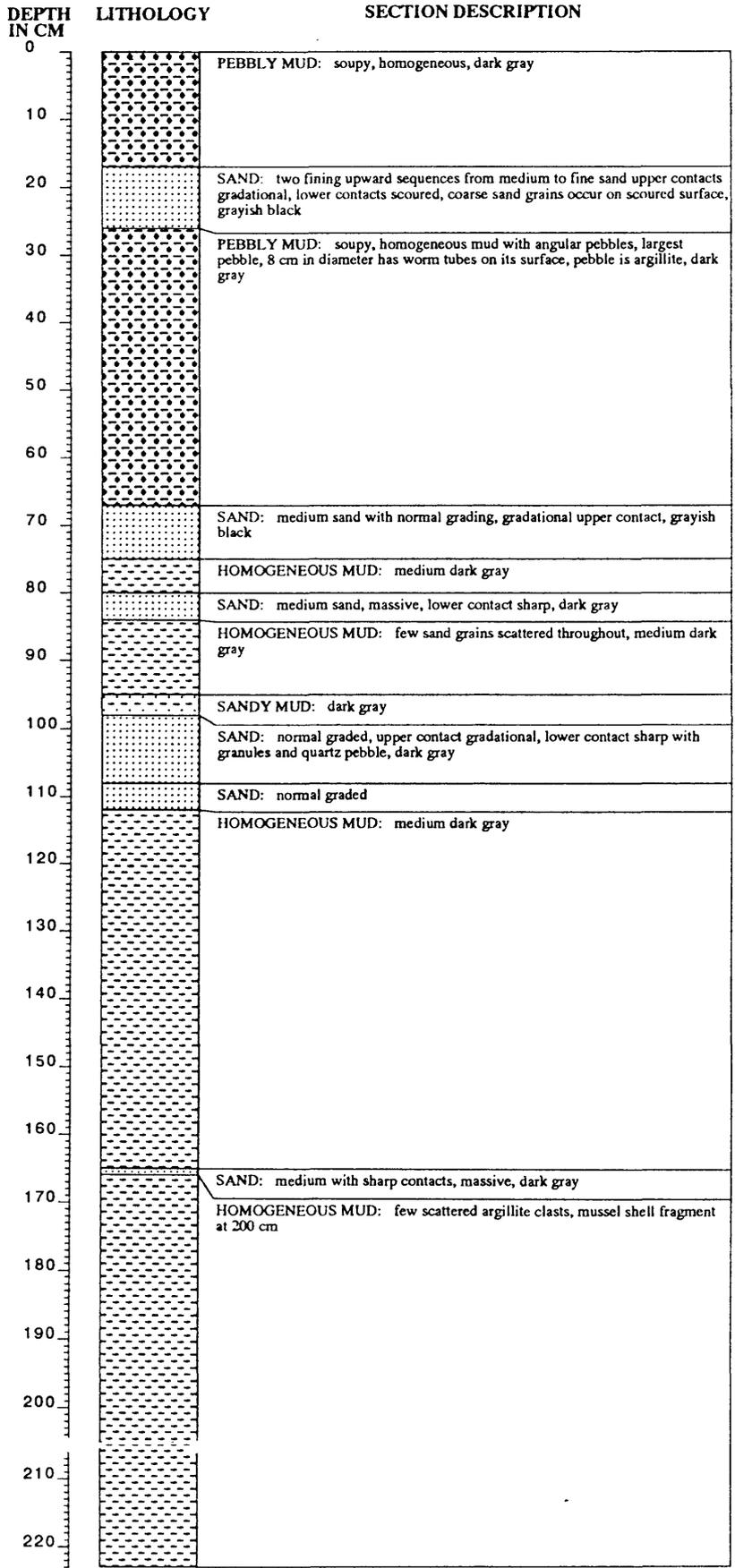
APPENDIX 1: CORE LOGS

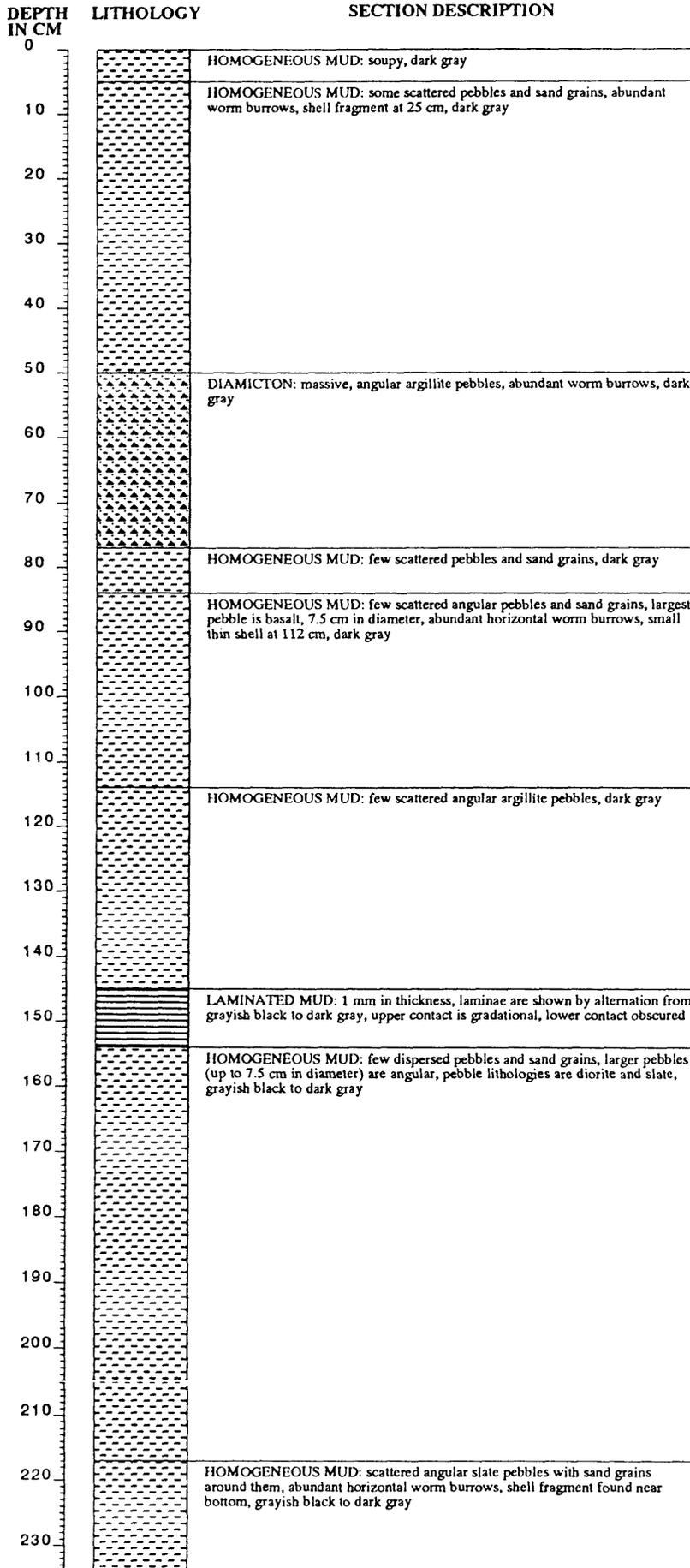
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		SANDY MUD: with large pebbles between 6-7 cm in diameter, dark gray
10		SAND: sharp contacts, grayish black
20		HOMOGENEOUS MUD: dark gray
		SILT: weakly laminated, fines upward, gradational upper contact, dark gray to grayish black
30		SANDY MUD: with large pebbles about 3 cm in diameter, dark gray
40		SILT: dark gray
50		HOMOGENEOUS MUD: with isolated small pebbles, dark gray
60		SAND: massive, sharp upper and lower contacts, grayish black
70		HOMOGENEOUS MUD: very soupy, dark gray
80		
90		
100		
110		
120		
130		
140		
150		
160		

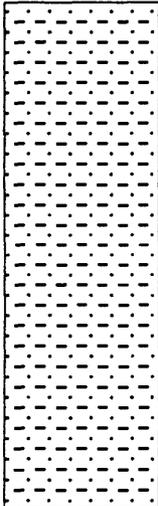
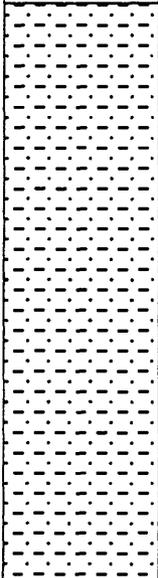
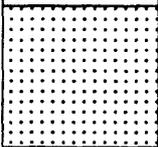
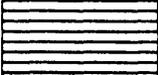


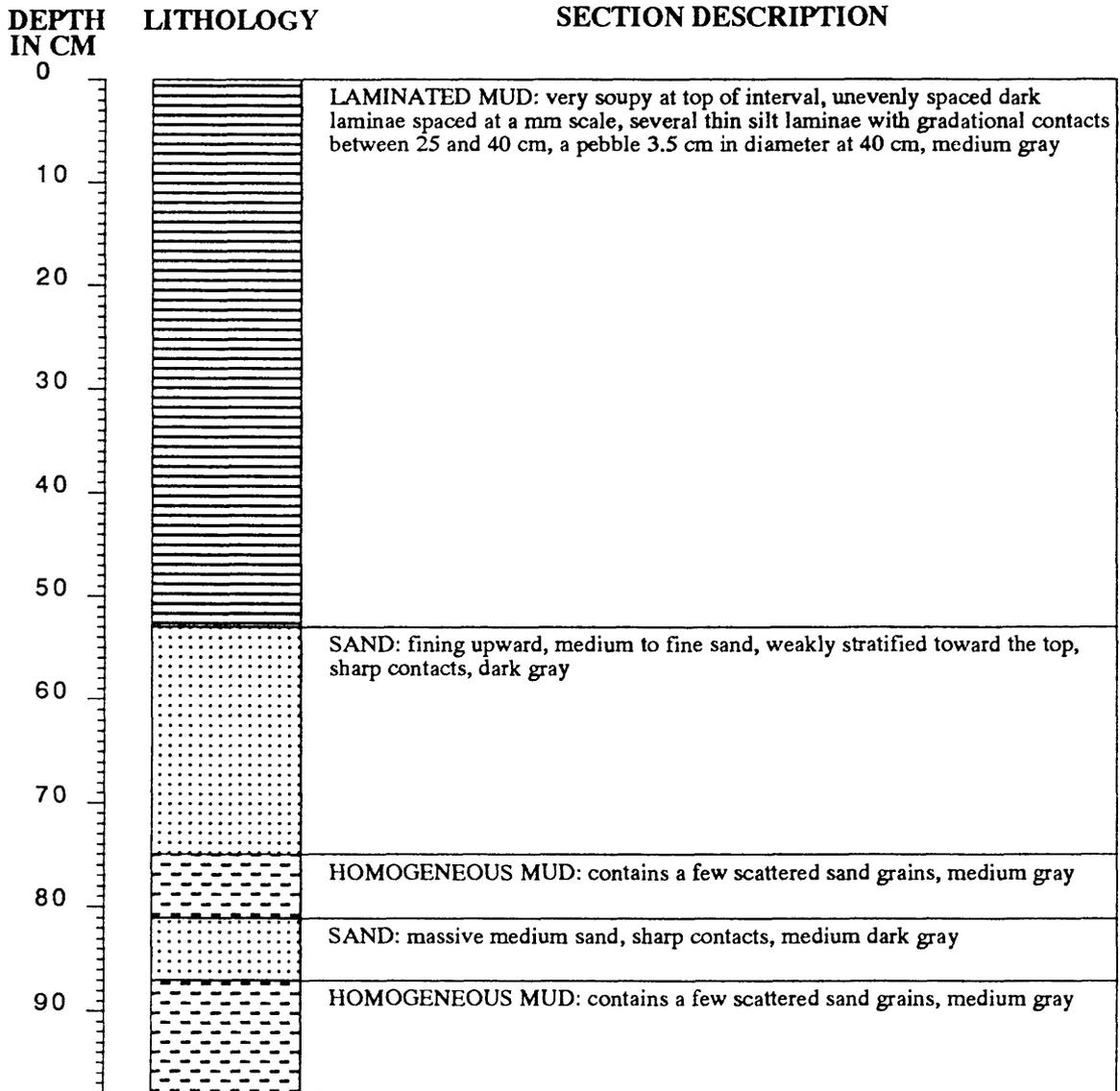
AH93 PW GC-3

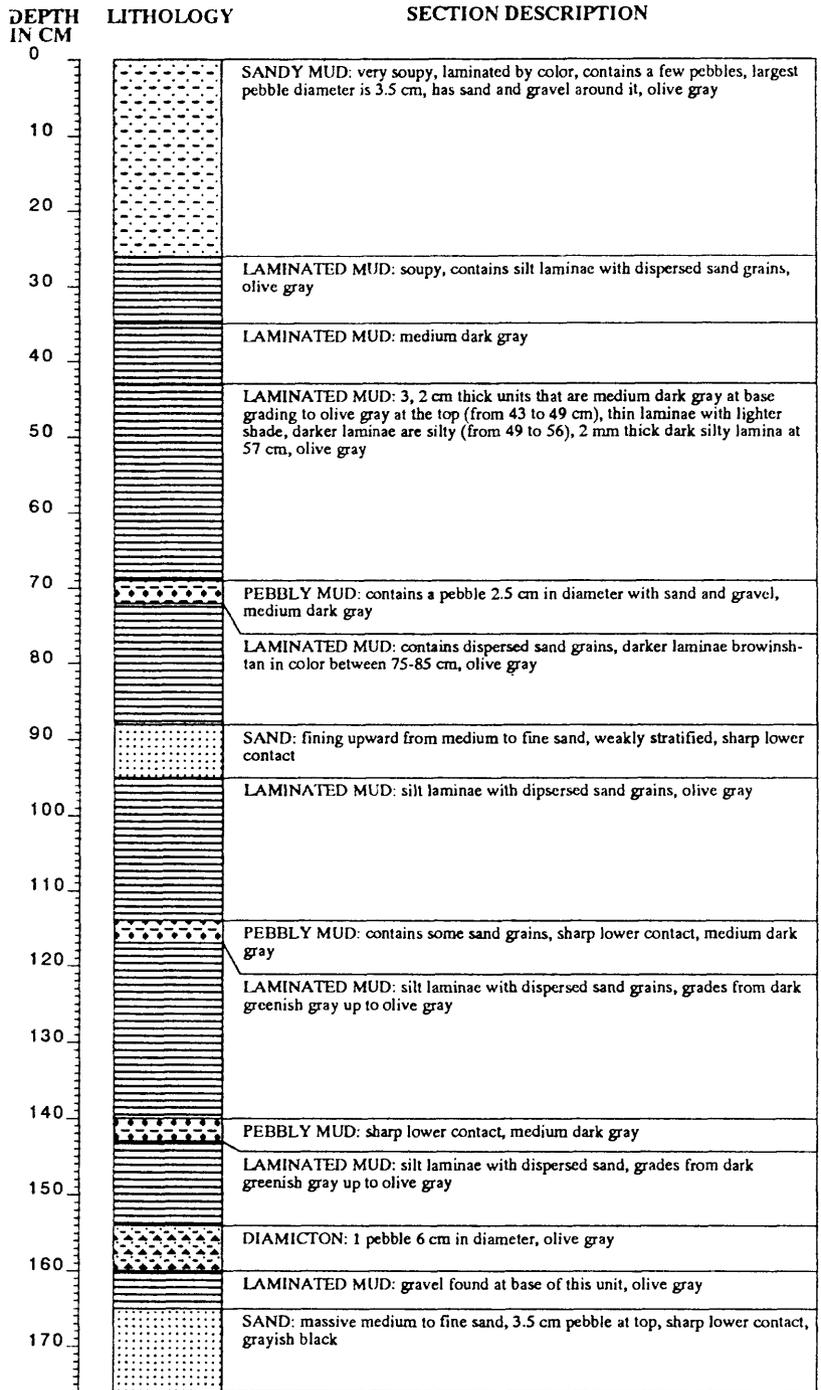
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		SANDY MUD: soupy, few scattered small clasts, dark gray
10		PEBBLY MUD: with sand, angular argillite pebbles, largest is 3.5 cm, lower contact is gradational, black worm tube 10 cm long, 2-3 mm in diameter, dark gray
20		DIAMICTON: higher percent sand and pebbles than above, largest angular clast is 4 cm, dark gray
30		
40		

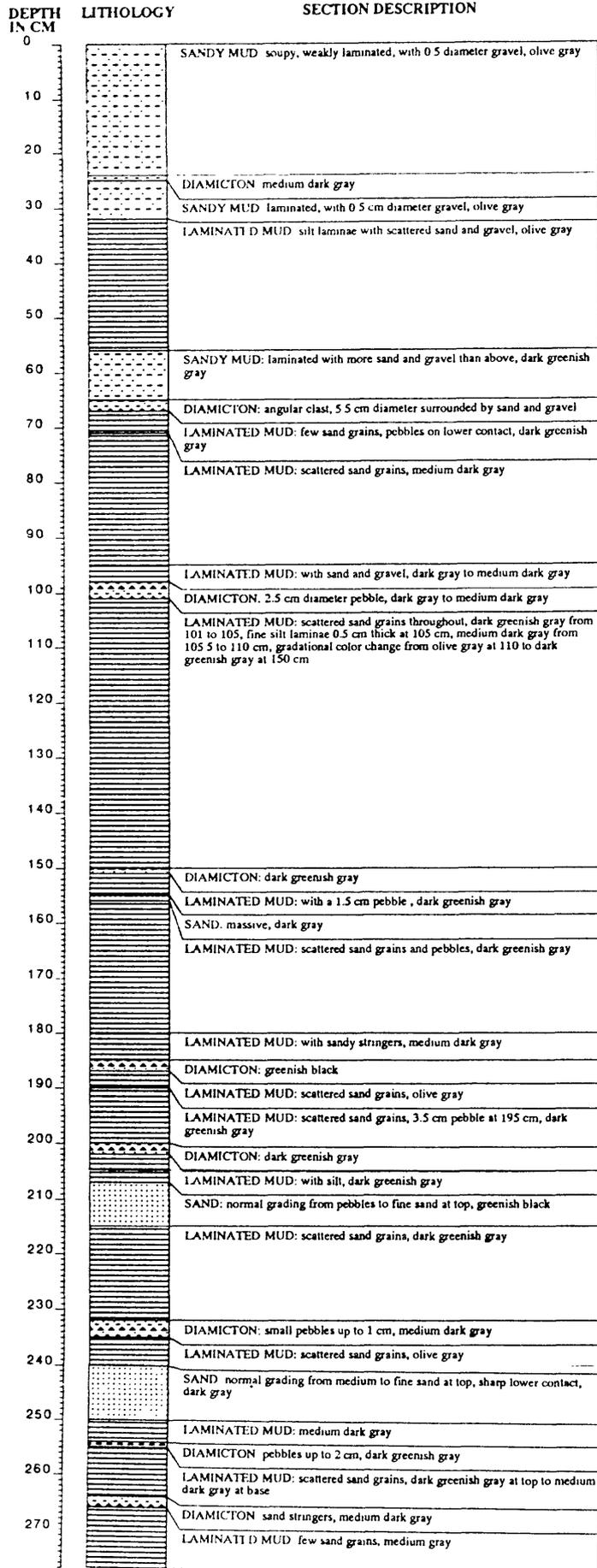


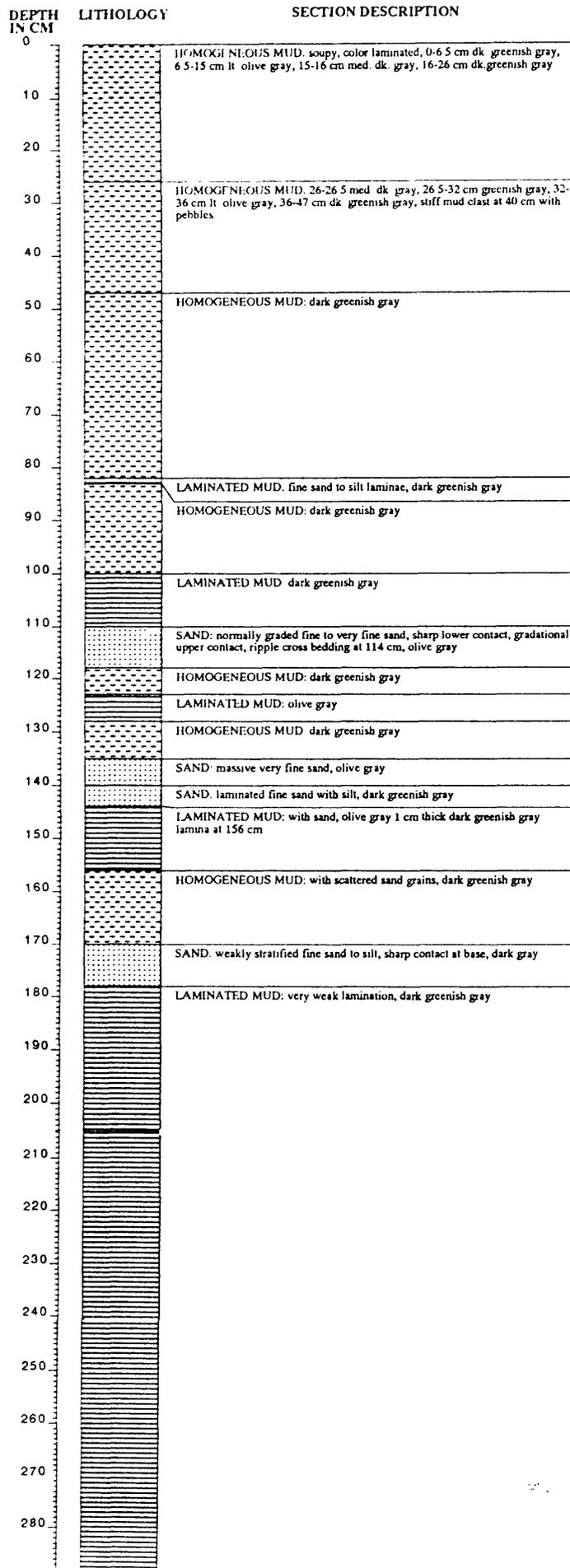


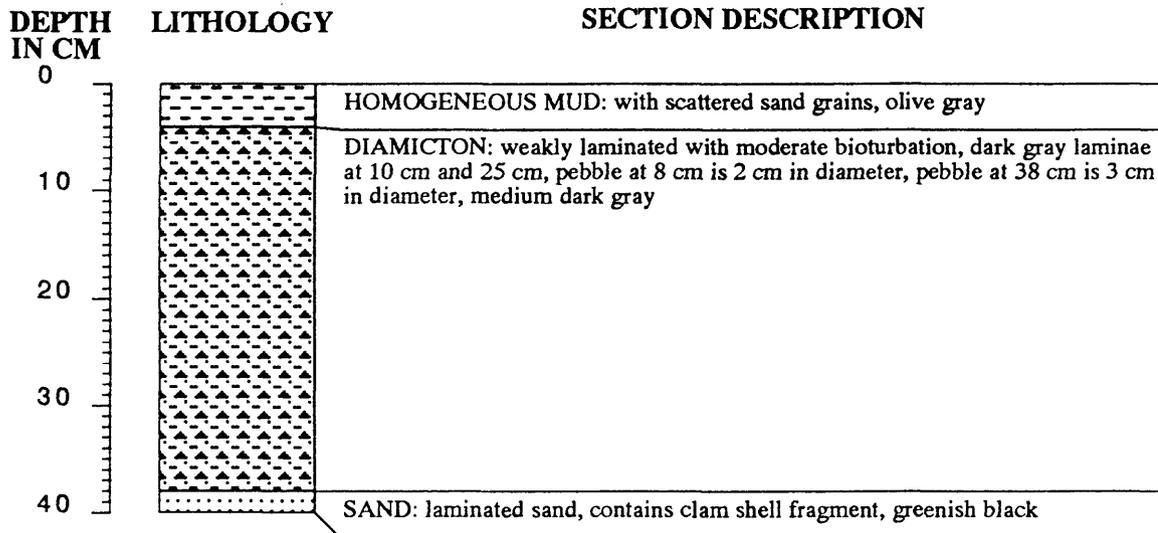
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0 10 20 30 40		SANDY MUD: very soupy and disturbed, medium dark gray
50 60 70 80 90 100		SANDY MUD: less soupy, not disturbed, medium dark gray
110		SAND: fine sand, normal graded, upper contact is gradational, lower contact is sharp, medium dark gray
120		LAMINATED MUD: 1-2 mm in thickness, scattered sand grains and pebbles up to 2 cm

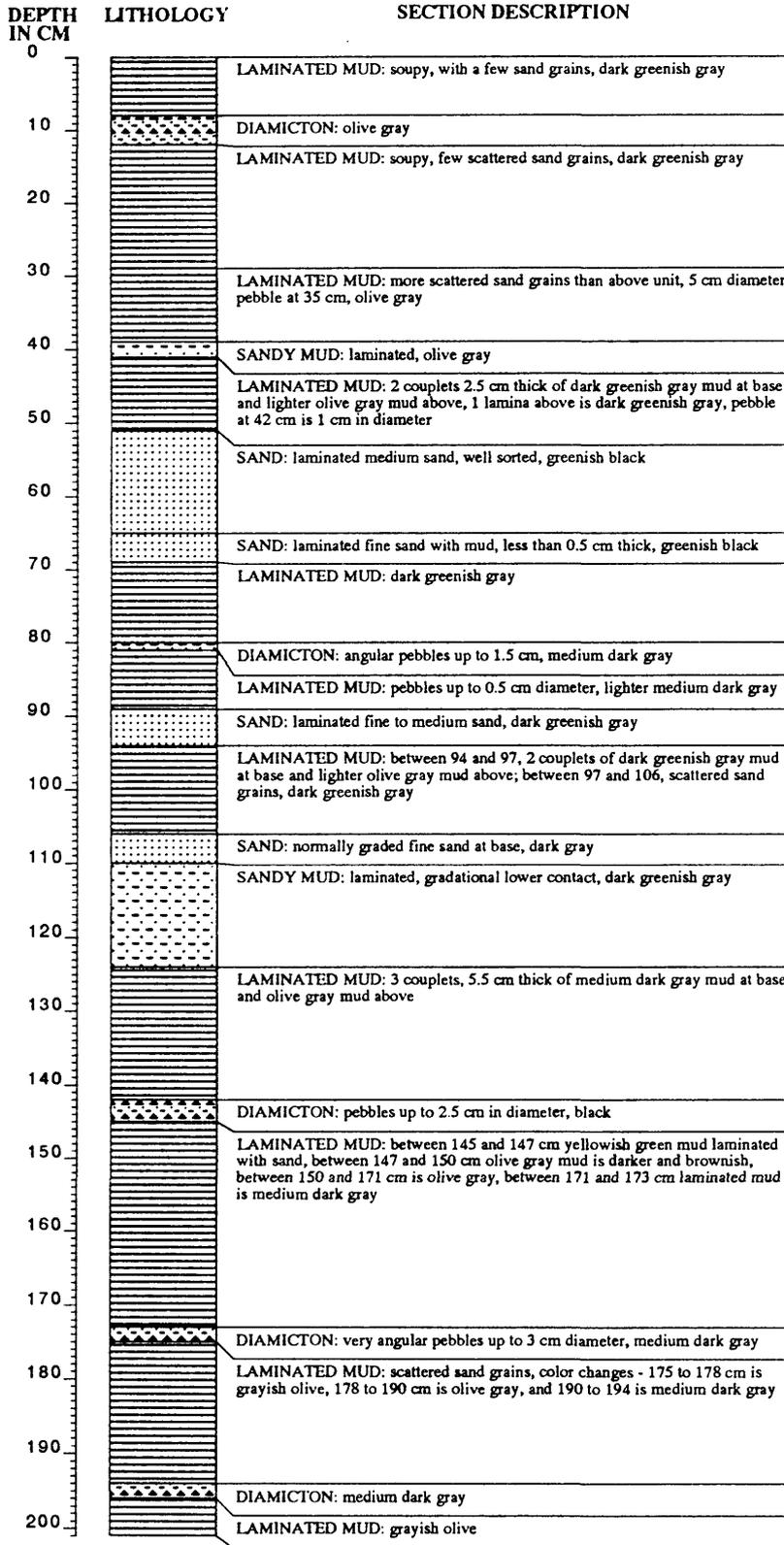


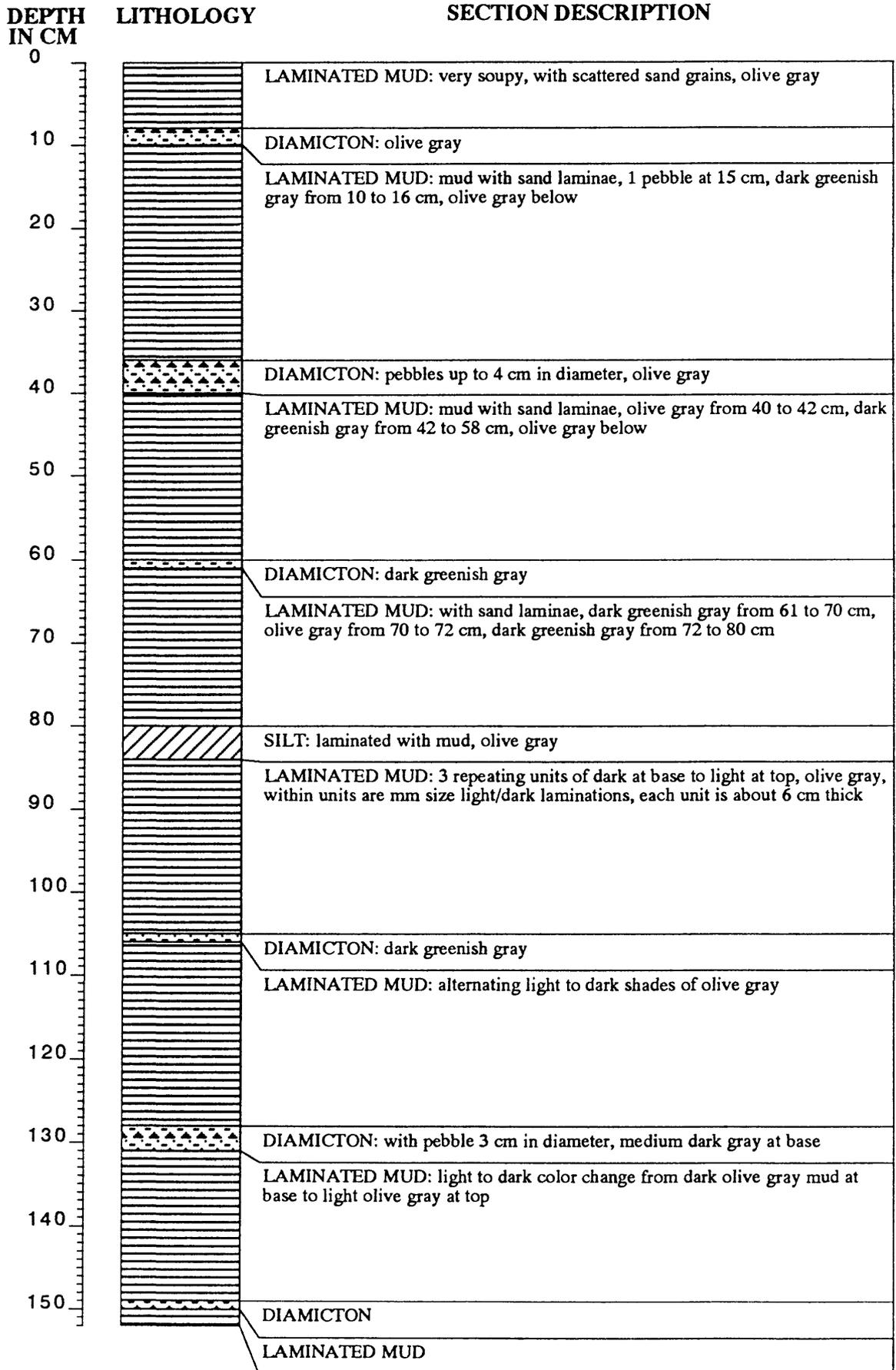


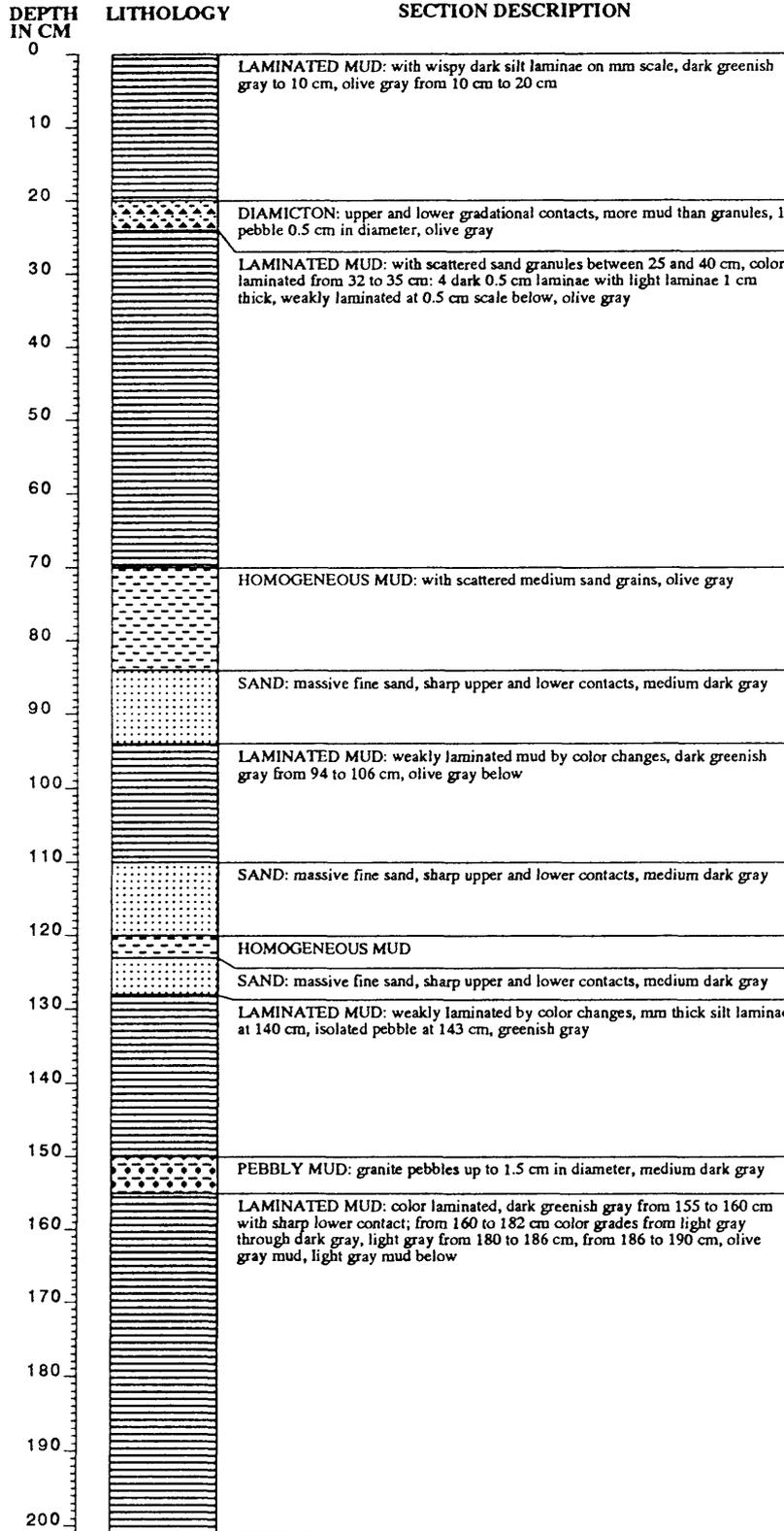


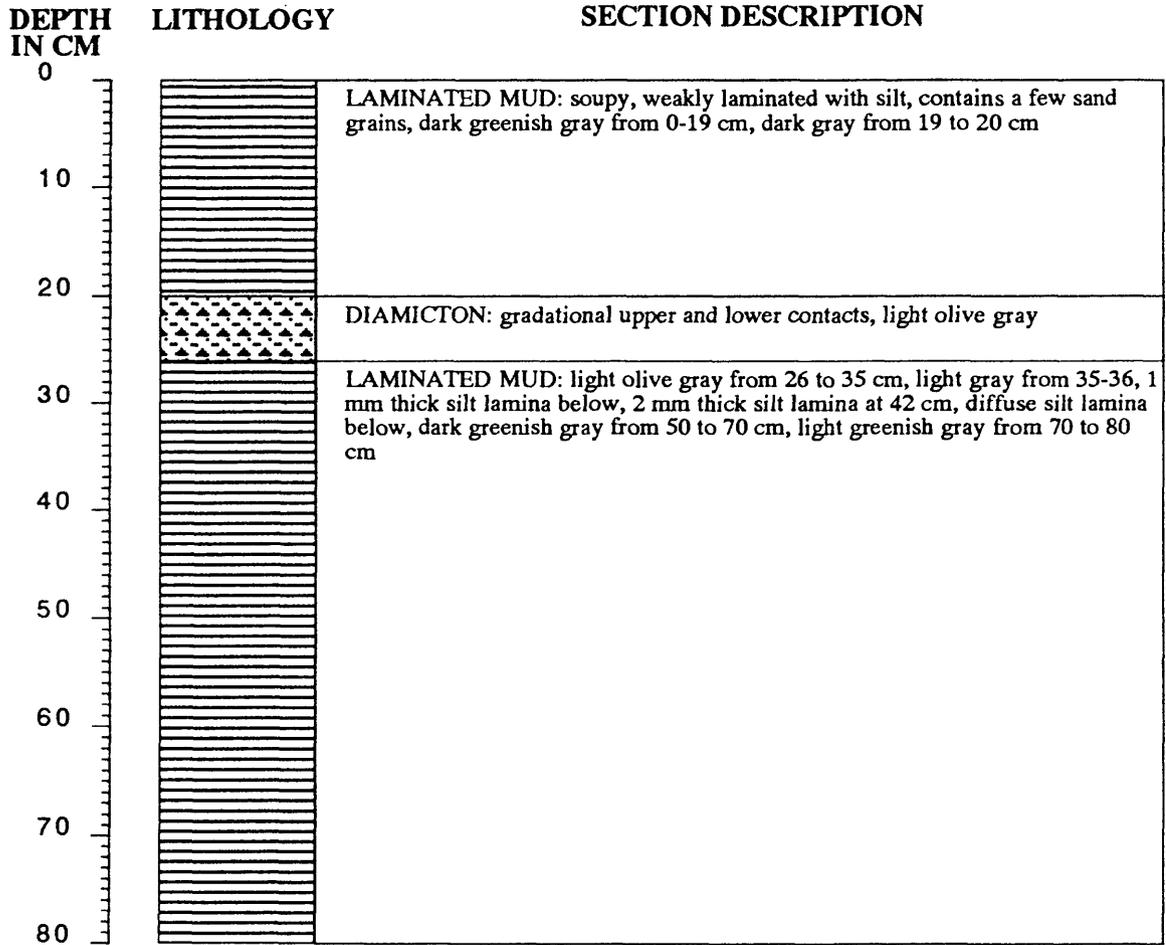


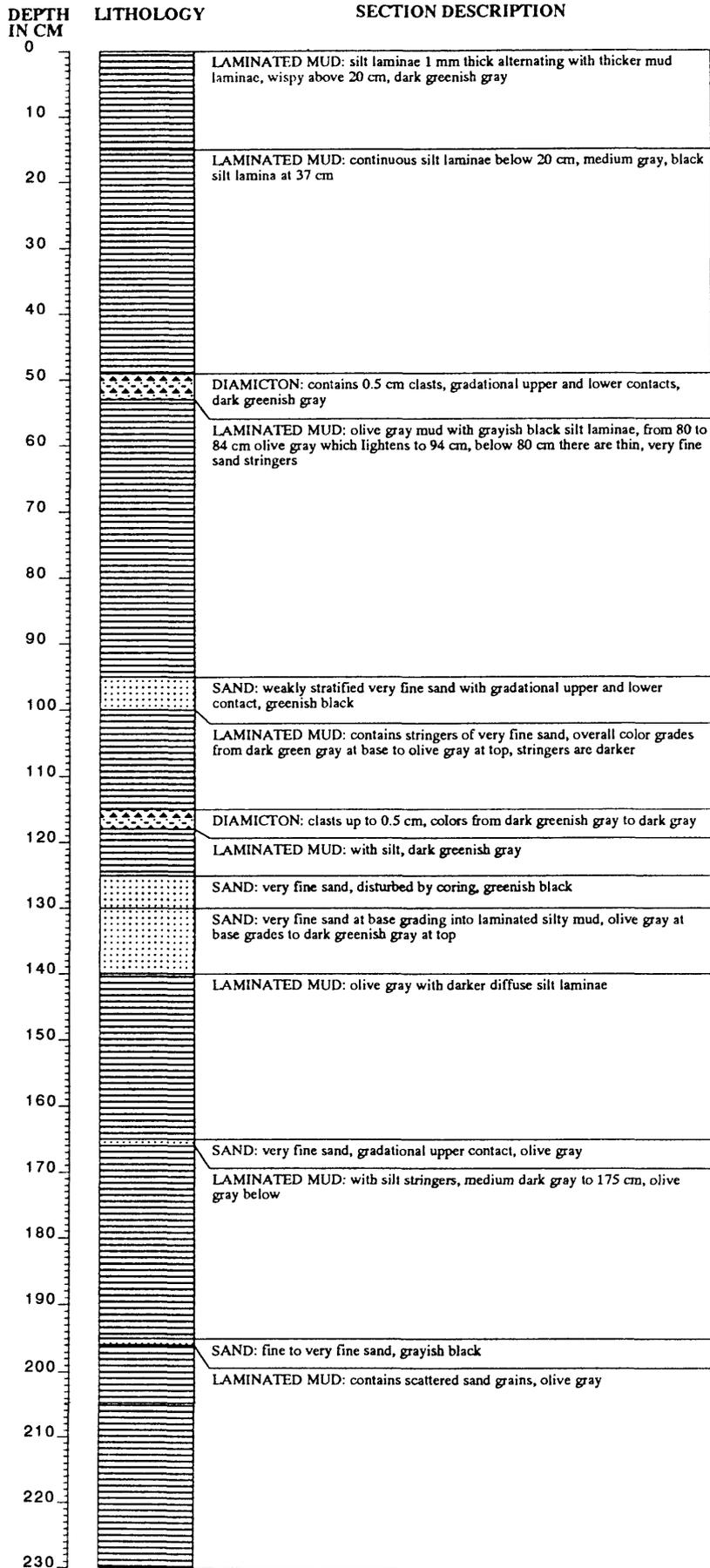


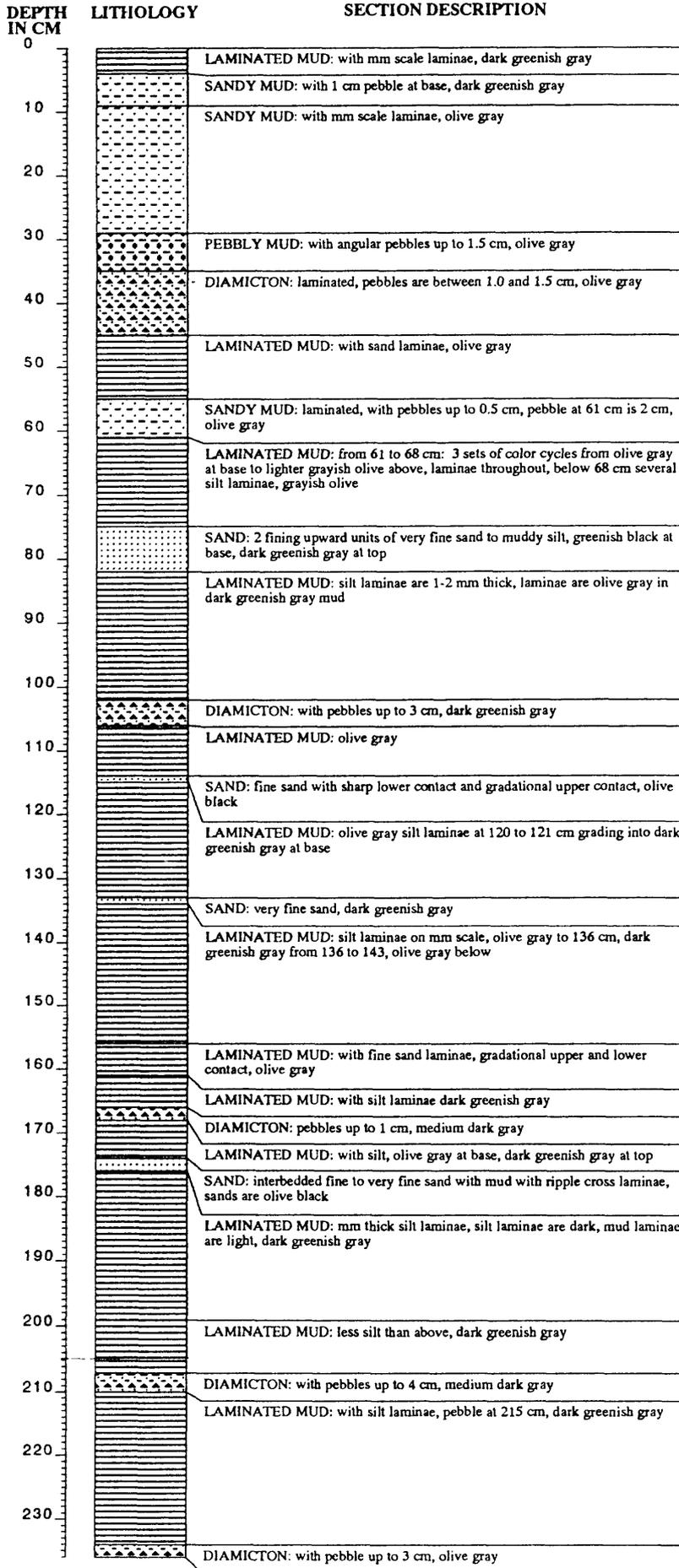


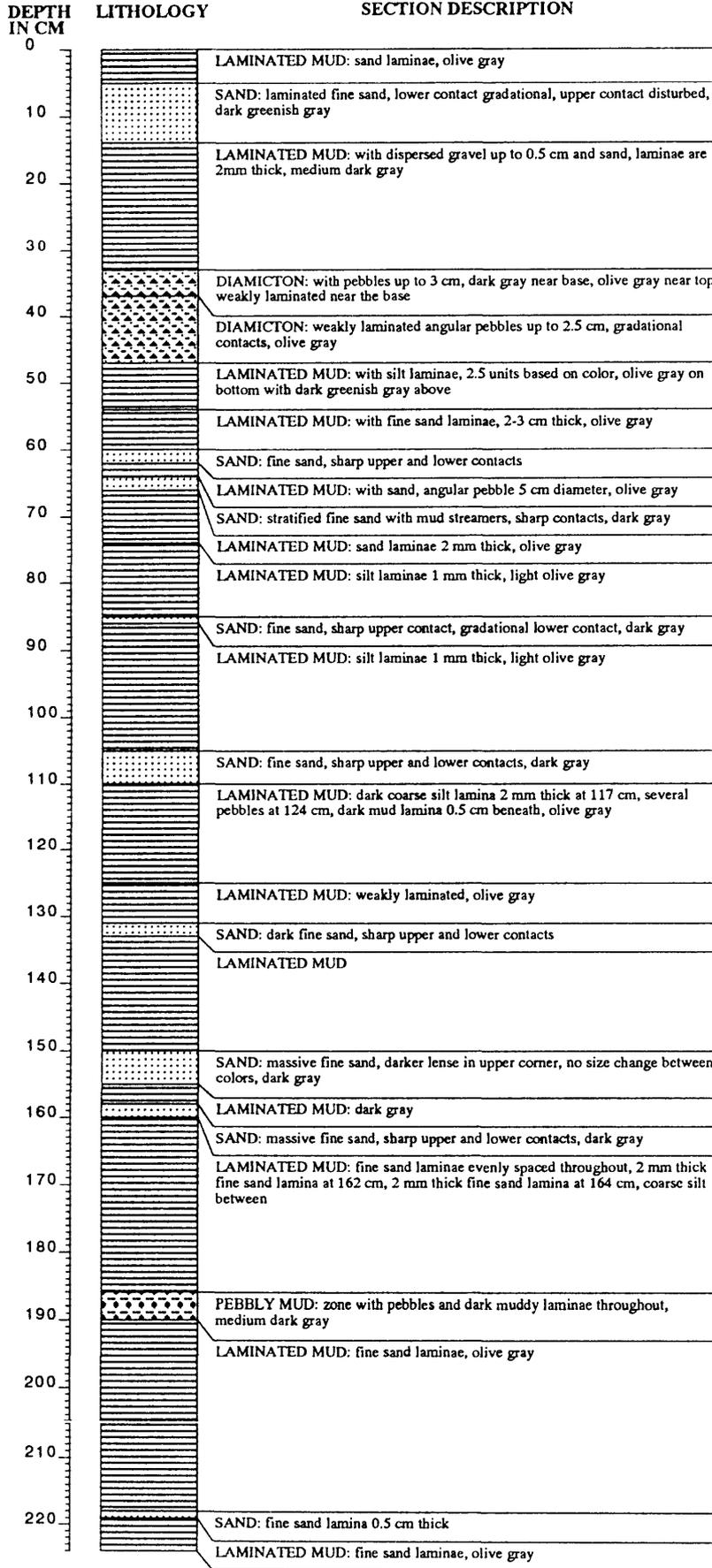


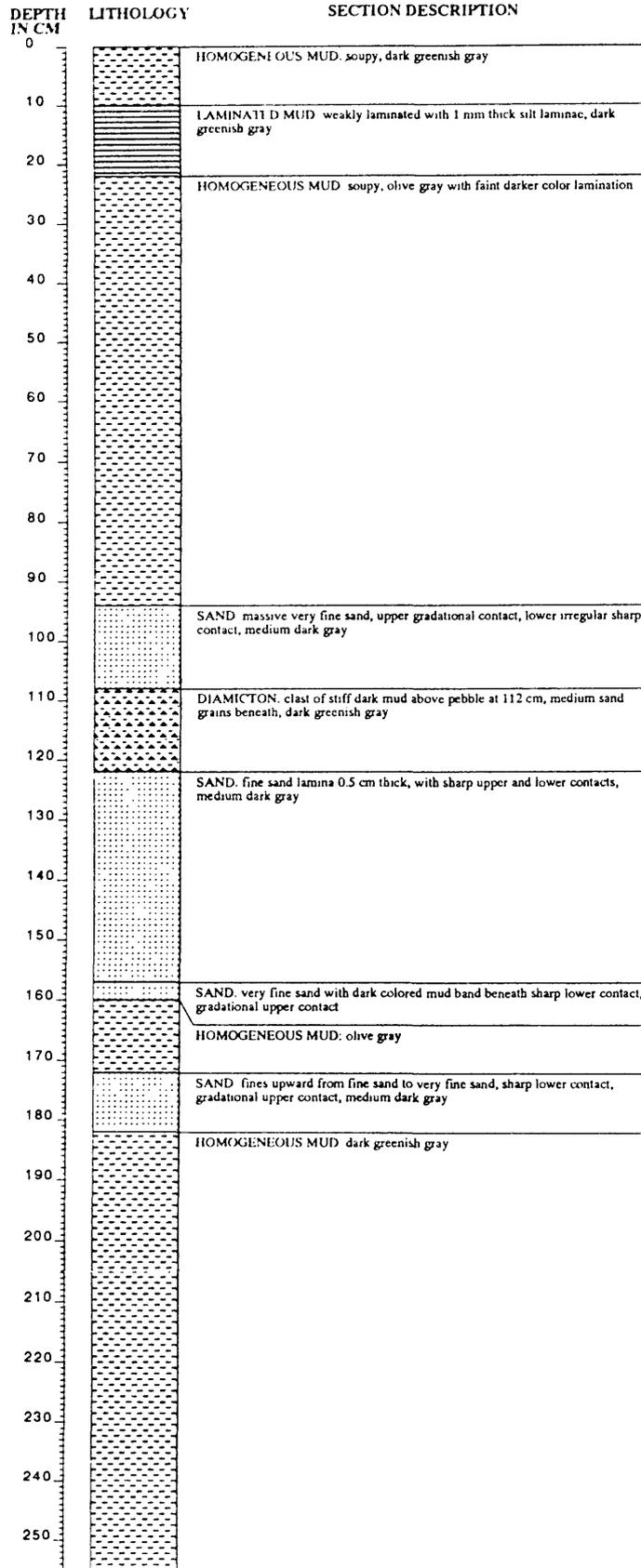


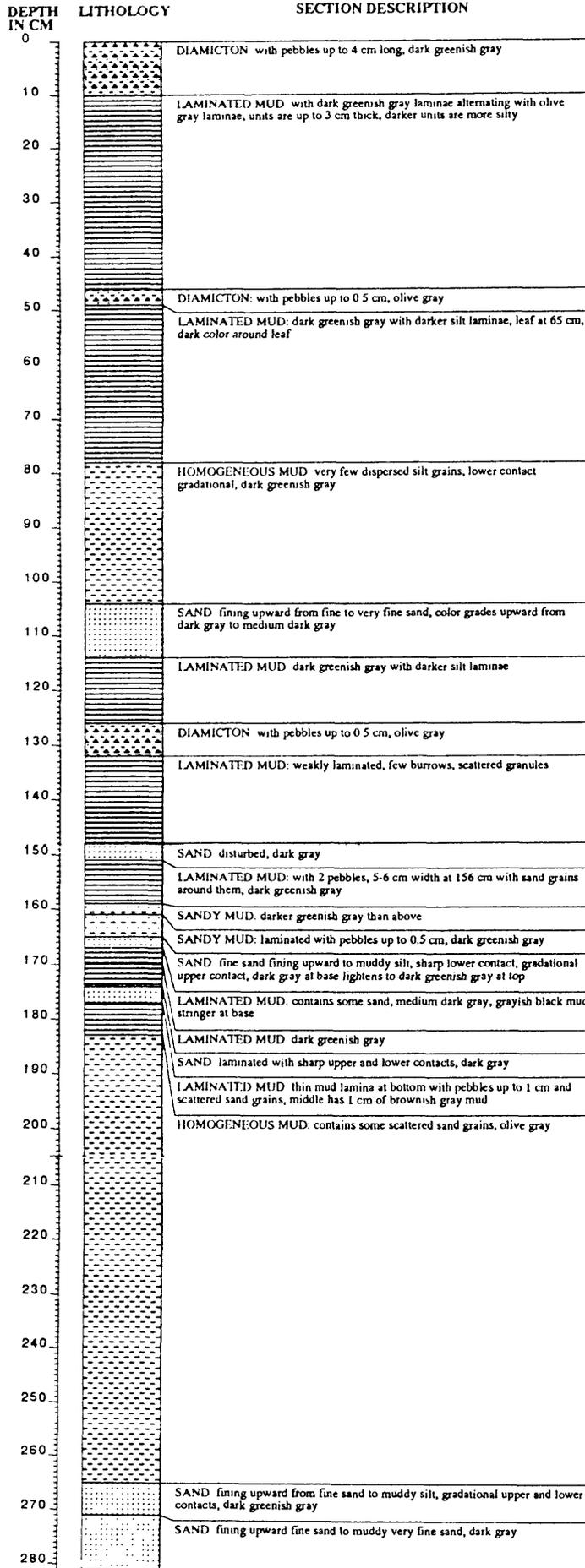


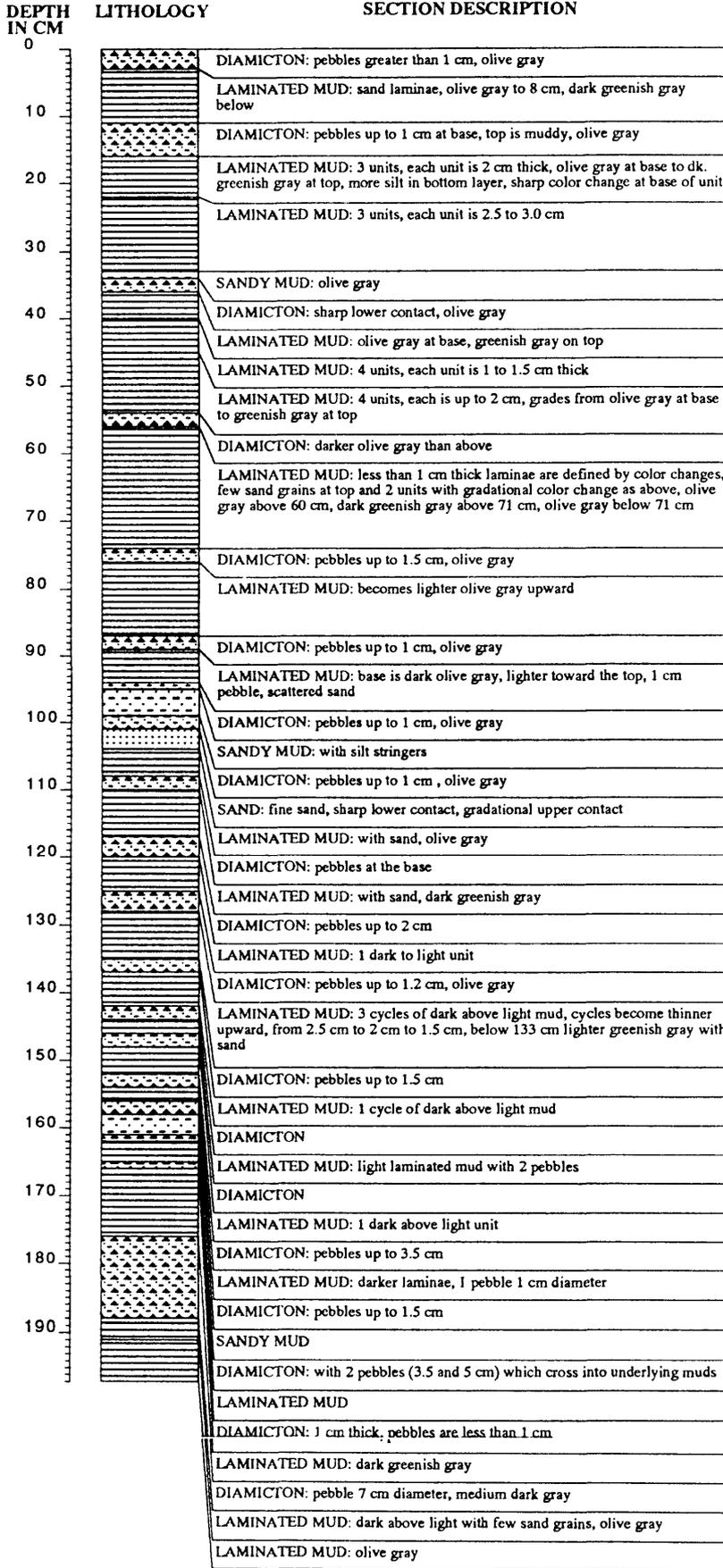


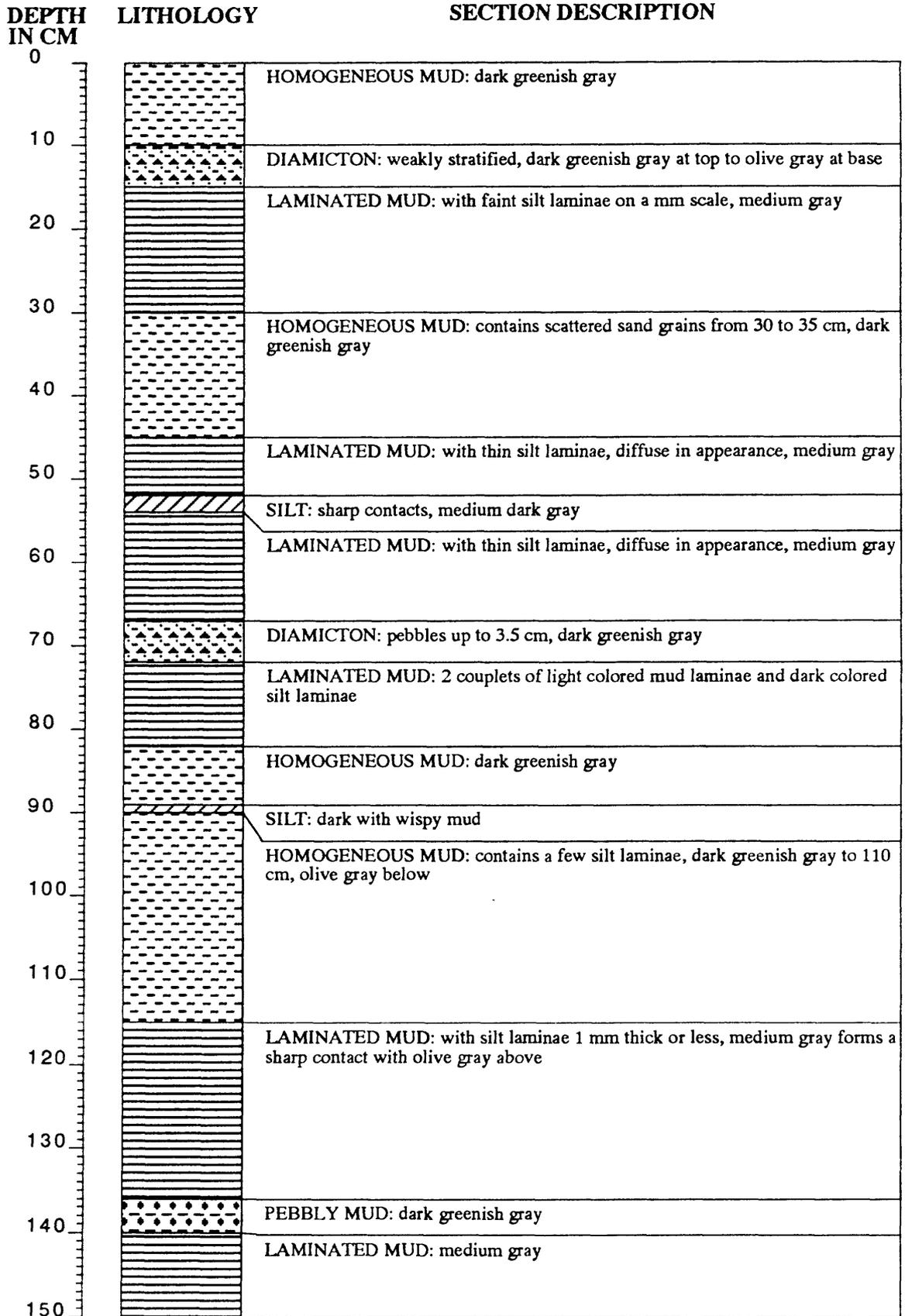


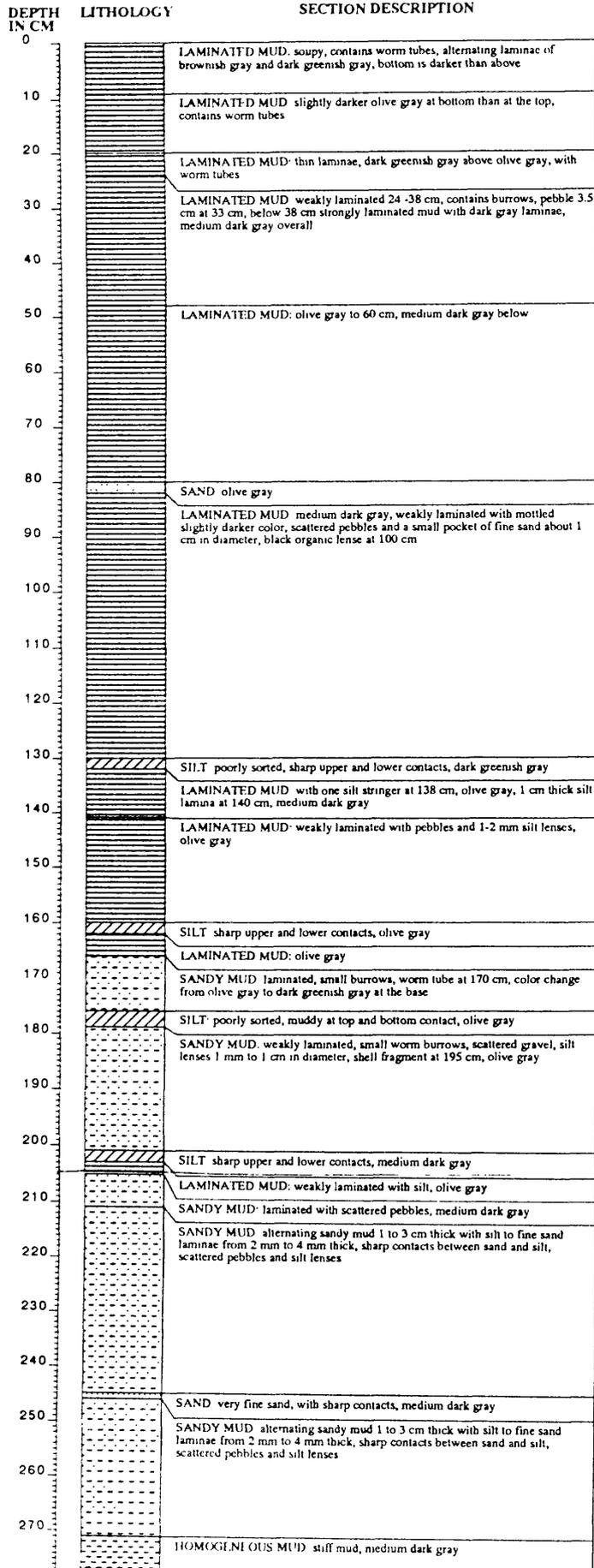




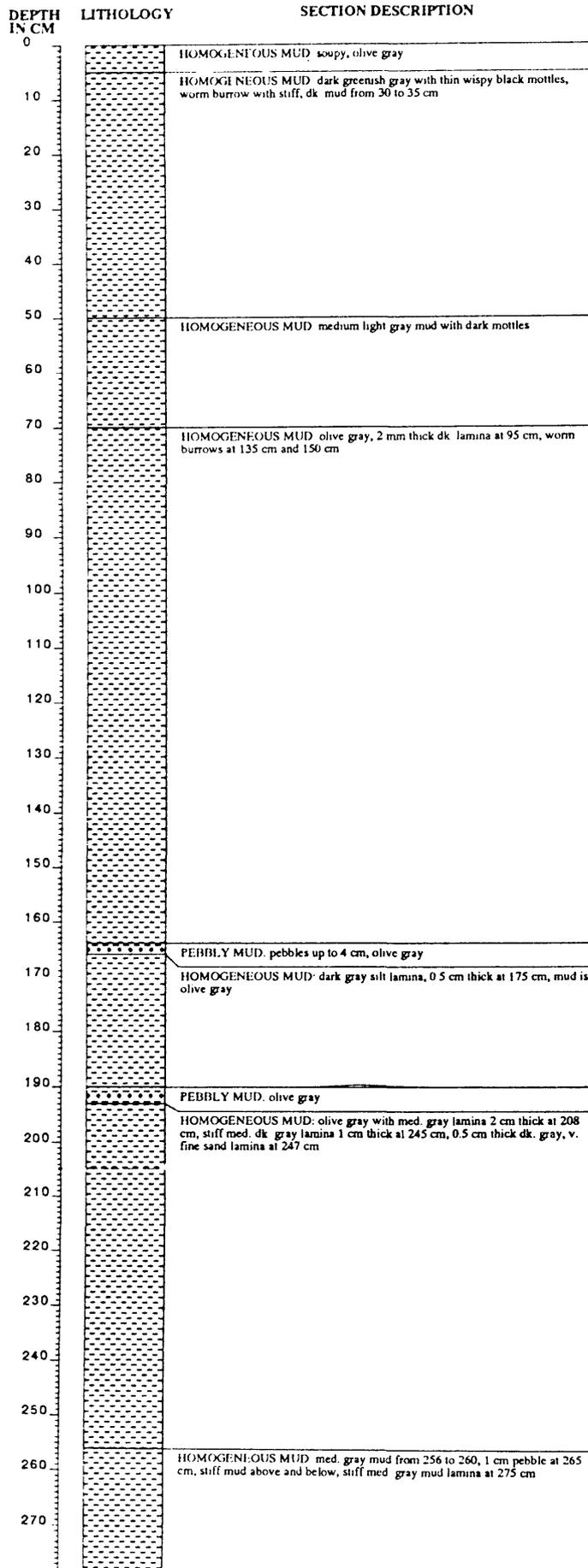


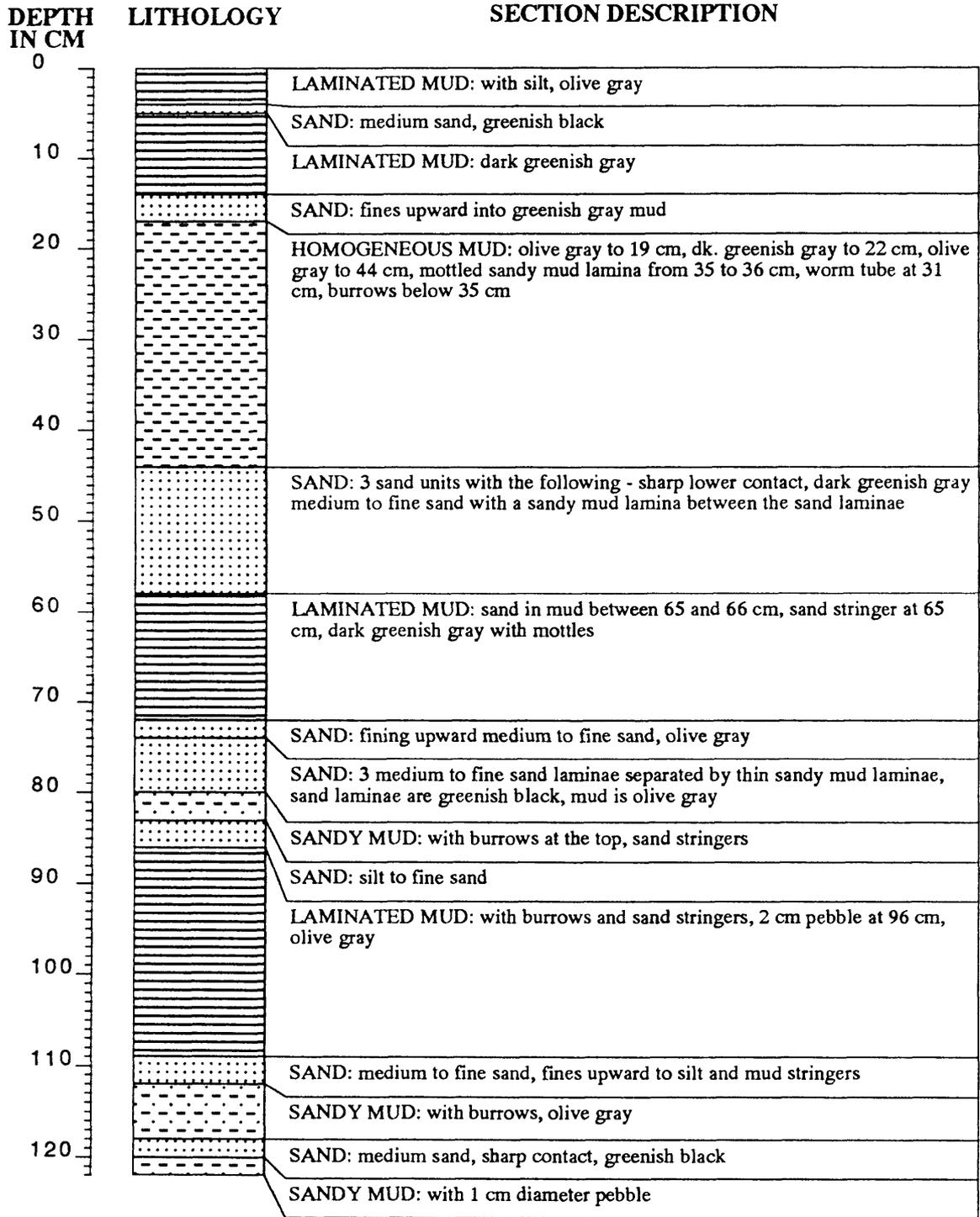


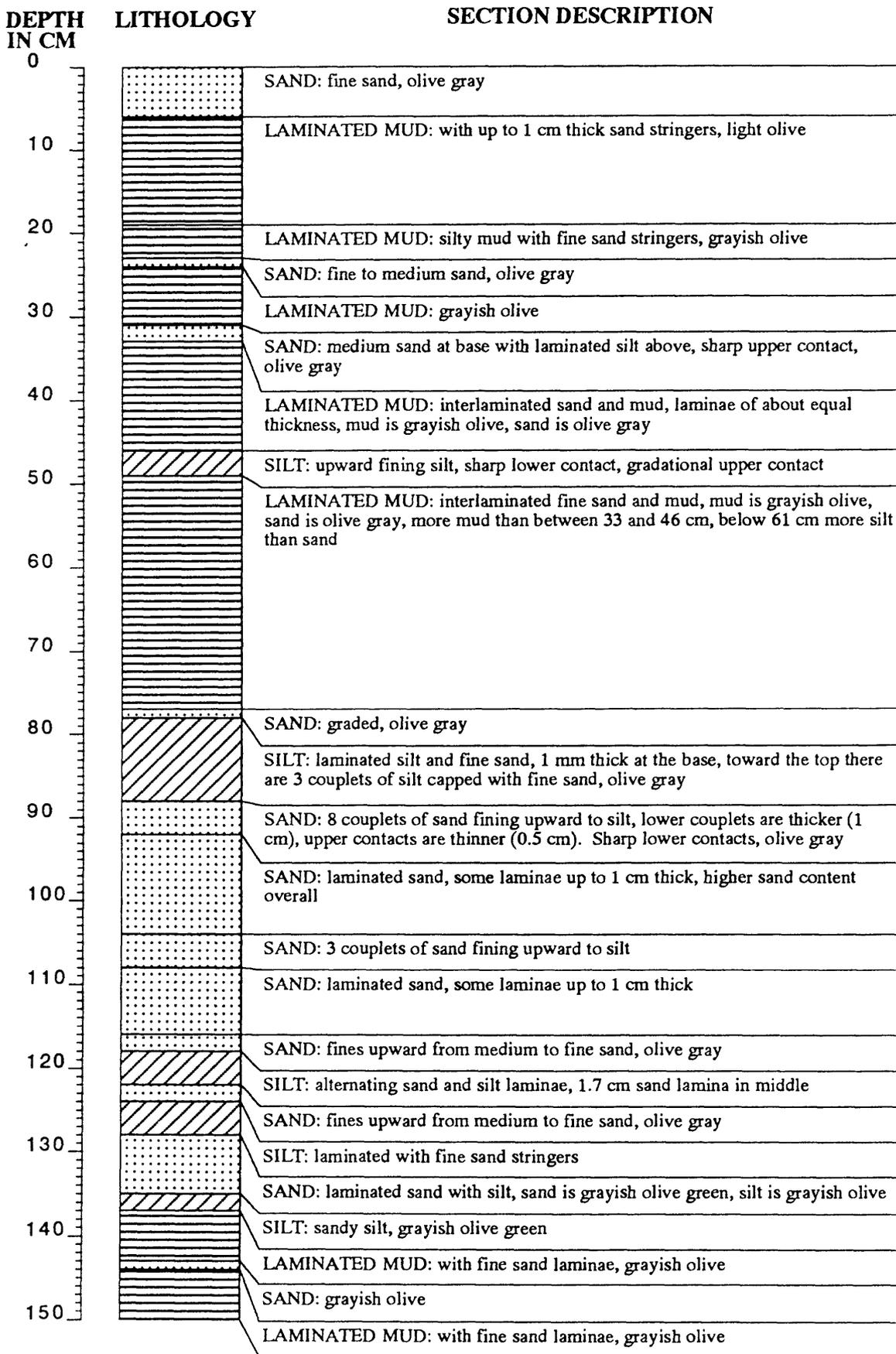


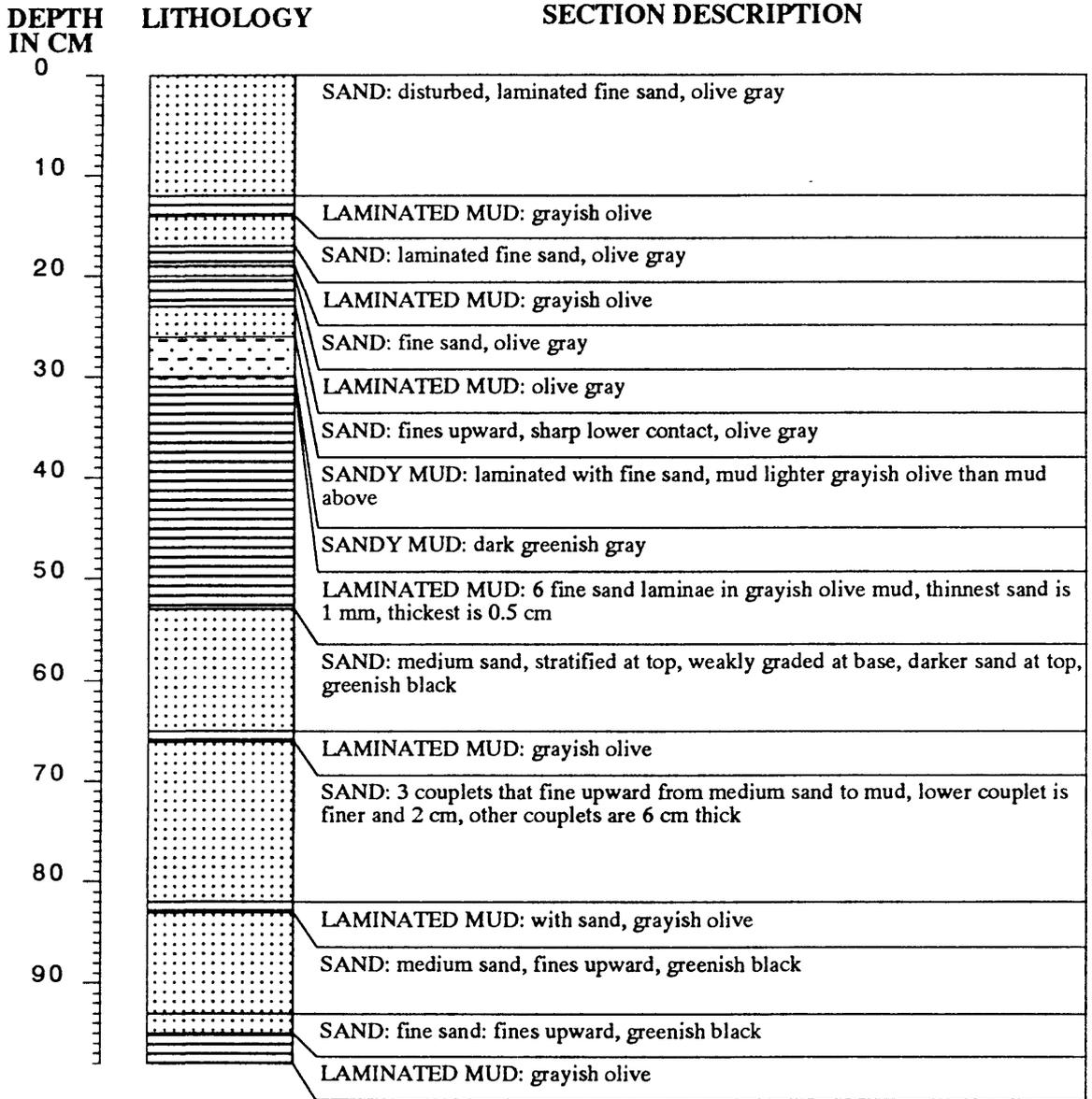


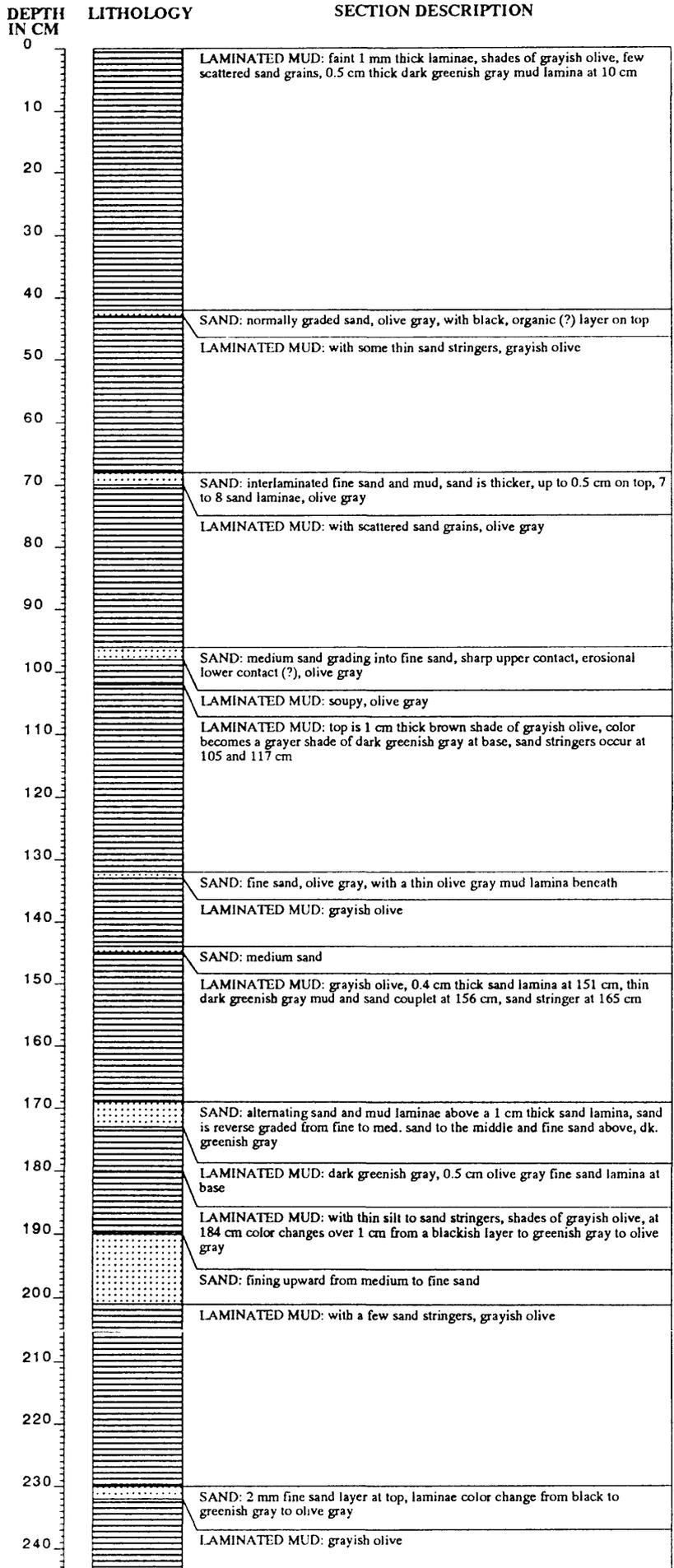
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		LAMINATED MUD worm tube at top, soupy, shades of olive gray
10		LAMINATED MUD from 10 to 20 cm grayish olive, light olive laminae 1 mm thick to 29 cm, olive gray below, silt stringer at 27 cm
20		
30		
40		
50		LAMINATED MUD with worm burrows, organic stringers with mottled texture, diffuse or gradational color changes between dk. greenish gray and olive gray, 1 cm thick sand lamina at 80 cm, organic layer at 95 cm, pebbles at 115 cm
60		
70		
80		
90		
100		
110		
120		
130		
140		LAMINATED MUD less organic material than above, pebbles at 150, 190, 230 cm, worm burrows filled with silt or organics are common, scattered sand grains
150		
160		
170		
180		
190		
200		
210		
220		
230		
240		LAMINATED MUD with more organic material than above, greenish black
250		SILT: 1 cm pebble at top, 2 silt clasts, 3.5 cm long and 1 cm wide, 1 cm long and 0.5 cm wide, silt clasts are cross laminated, olive gray
260		LAMINATED MUD with burrows at 258 cm and scattered sand grains, olive gray
270		SILT fines upward to clay, olive gray
280		

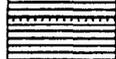
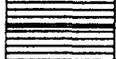


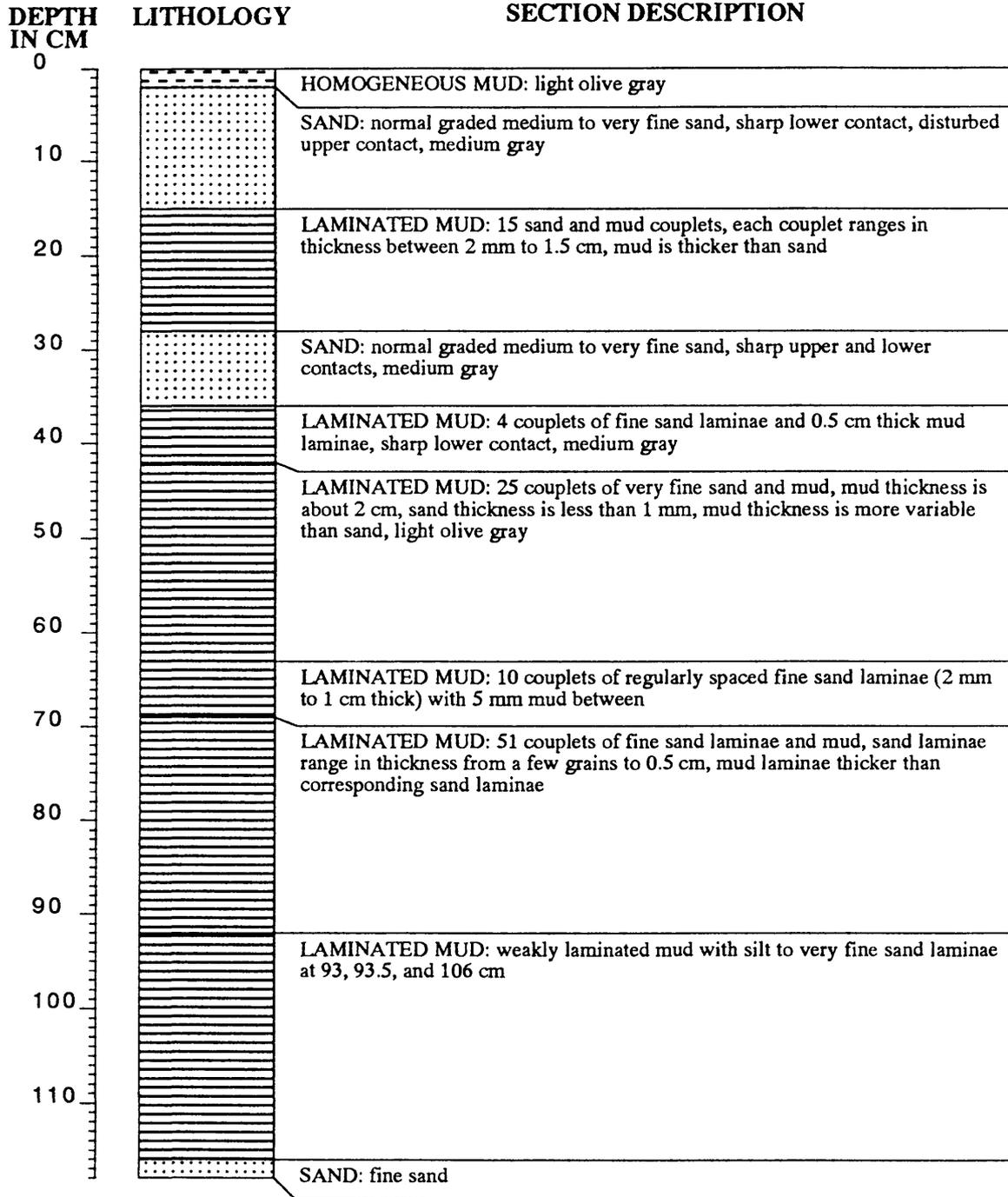


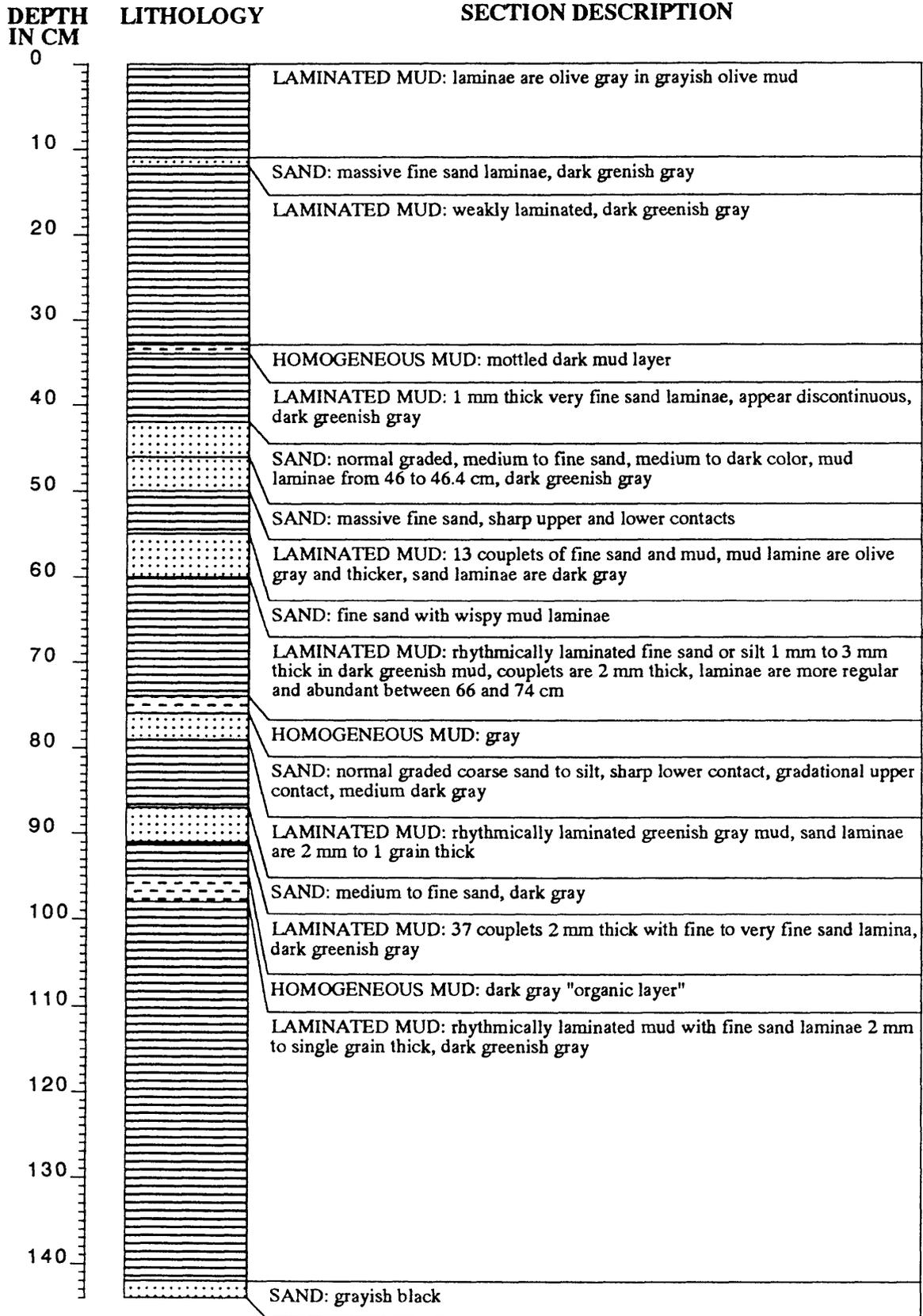




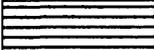
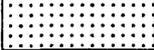
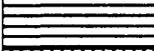
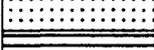
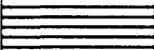
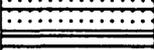
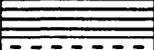
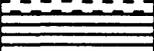
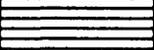
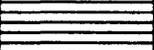


DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		LAMINATED MUD: dark laminae mm thick, greenish gray
10		LAMINATED MUD: 6 sand laminae 2-3 mm thick at 1.5 to 2.0 cm spacing from 20 to 32 cm, light olive gray
20		
30		
40		LAMINATED MUD: thick sand laminae above thin laminae, between 32 to 38 cm there are 8, 0.2-0.3 cm sand laminae, gradational upper and lower contacts, 40 to 46 cm fine silt laminae, from 46 to 50 cm there are 5, 2-5 mm thick sand laminae, greenish gray
50		LAMINATED MUD: weakly laminated
60		HOMOGENEOUS MUD: light olive gray
70		LAMINATED MUD: weakly laminated with silt, light olive gray
80		SAND: normal graded from coarse sand to silt, medium dark gray
90		LAMINATED MUD: from 70 to 76 cm light olive gray with brown and black 1 mm thick laminae, pebble at 72 cm, below 76 cm faintly laminated mud, greenish gray
100		LAMINATED MUD: silty laminae that range from 1 to 5 mm thick, dark greenish gray
110		LAMINATED MUD: sand laminae from 2 to 4 mm thick, olive gray
120		
130		SAND: fine sand, sharp contacts above and below, muddy lamina 3 mm thick at 128 cm, very fine sand from 128 to 131 cm
140		LAMINATED MUD: alternating olive gray and greenish gray, 2mm thick very fine sand lamina at 139 cm, greenish gray mud with 1 mm thick silt laminae
150		
160		SAND: fine sand lamina 0.5 cm
170		LAMINATED MUD: olive gray mud with faint laminae and dispersed sand grains
180		
190		
200		LAMINATED MUD: color laminae in shades of olive gray with a 6 mm black lamina at 229 cm, thicker sand laminae occur at 195, 197, 205, 210, 212, 216, 225, 232, 235, 238 cm
210		
220		
230		
240		

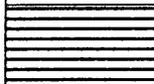
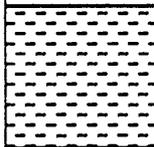
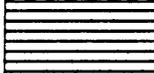
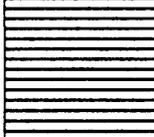
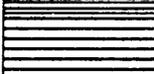
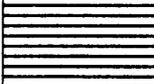
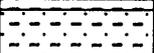
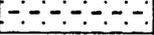


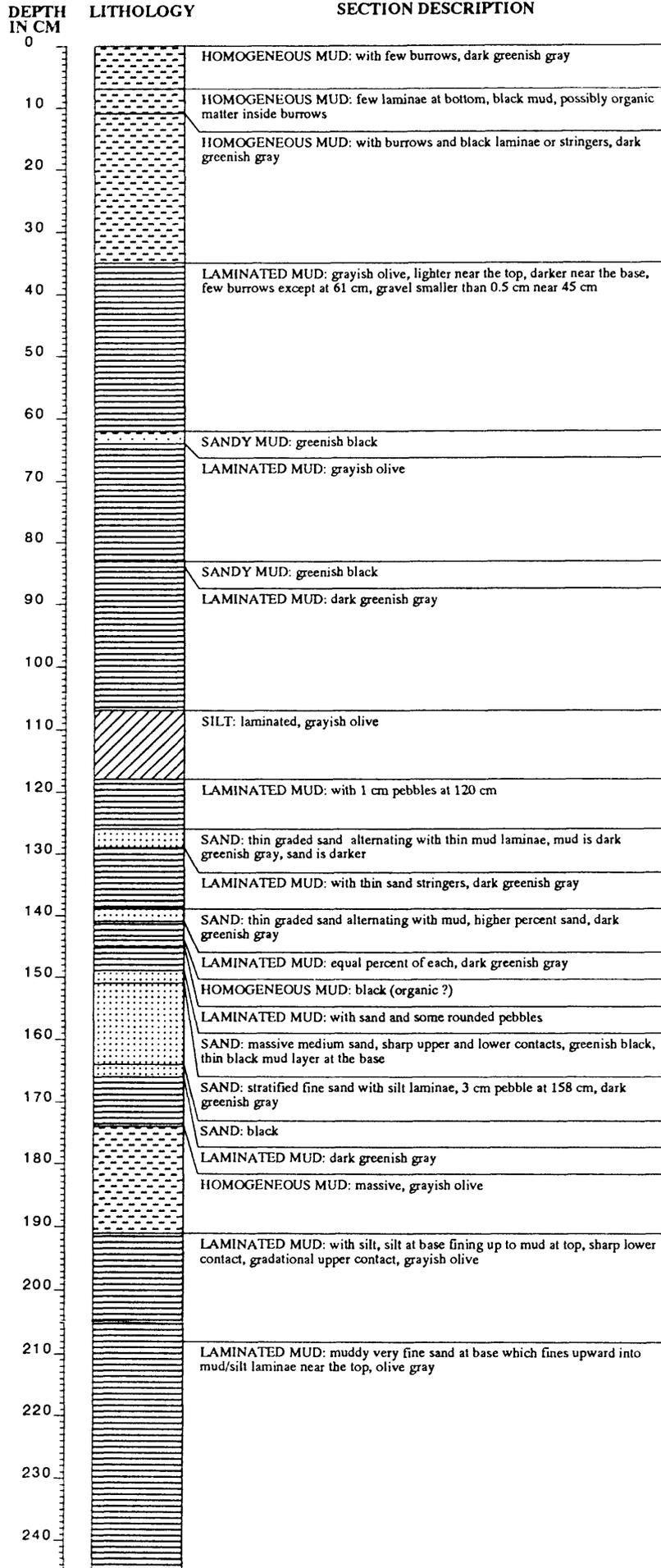


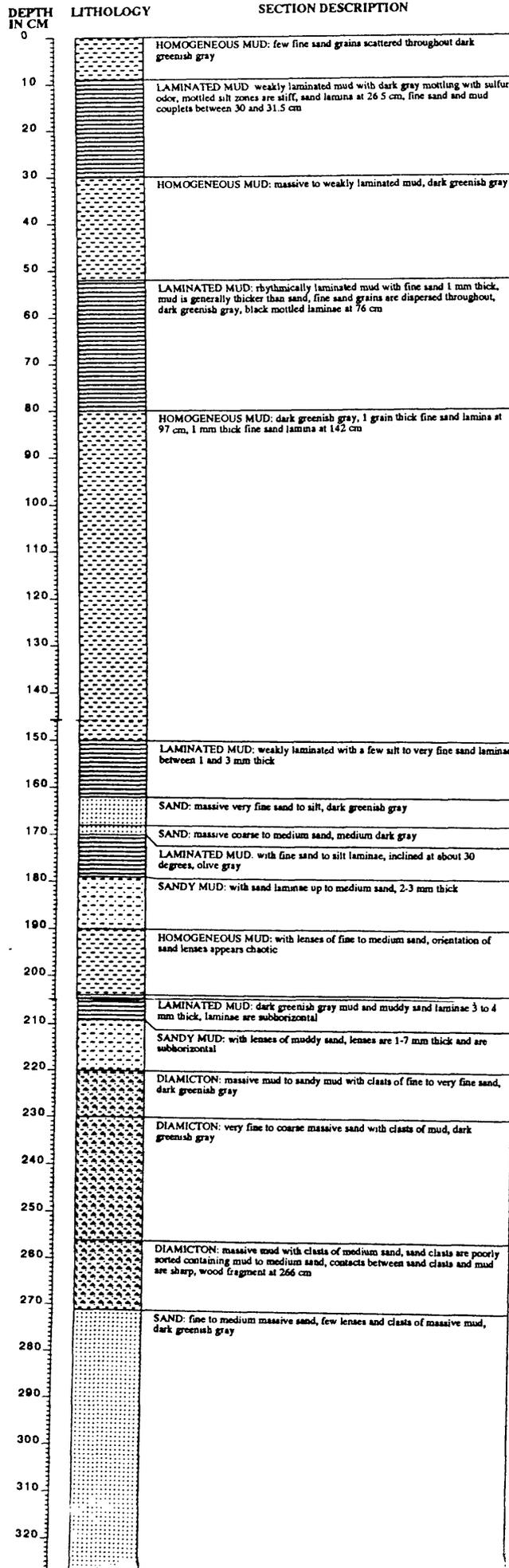
AH93 GB GC-10

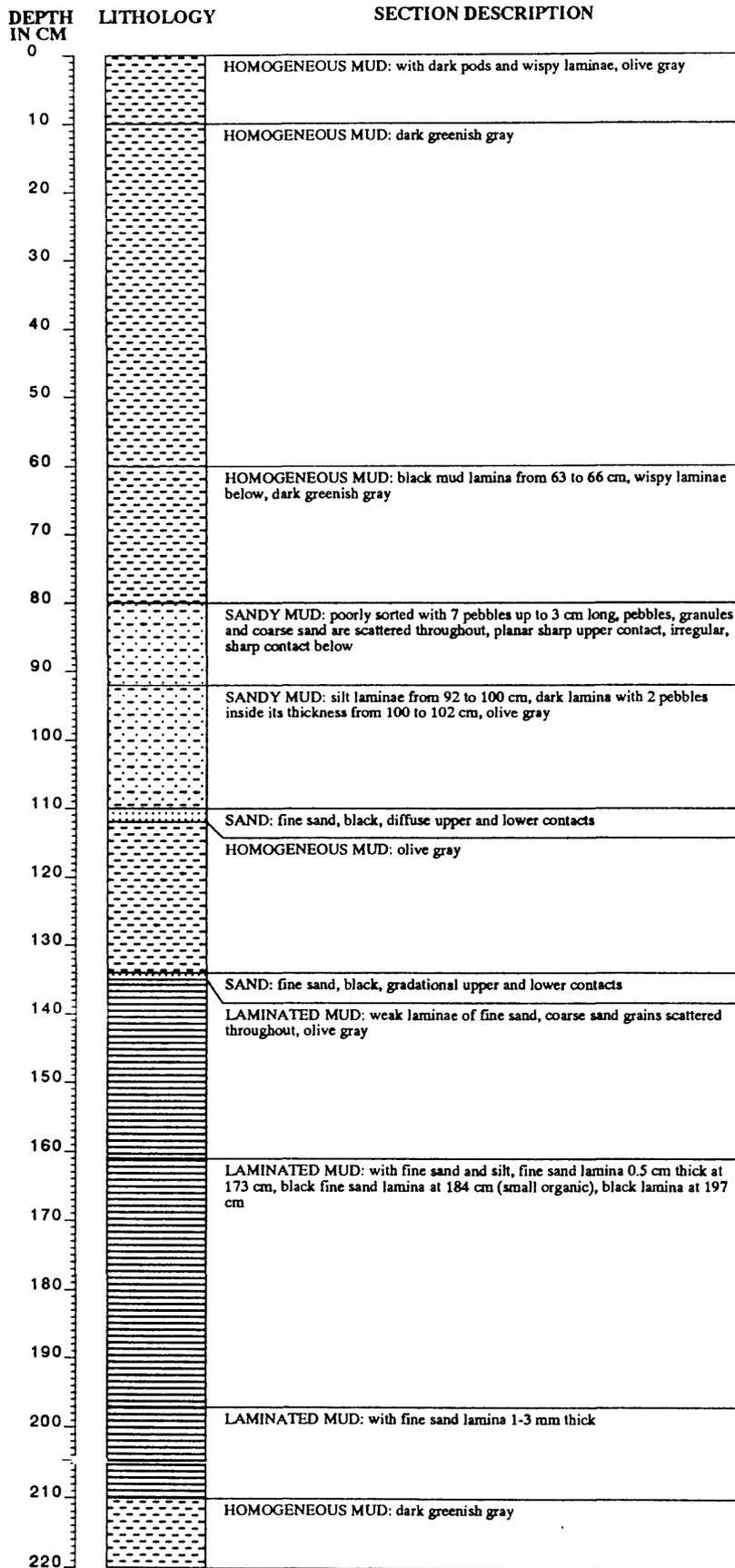
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		LAMINATED MUD: grayish olive
10		SAND: fines upward from medium to very fine sand, scoured base
		LAMINATED MUD: grayish olive
20		SAND: fines upward from medium to very fine sand, gradational upper contact
		LAMINATED MUD: grayish olive
40		SAND: fine to very fine sand grading upward into mud, dark greenish gray
		LAMINATED MUD: sand laminae, 38 to 41 cm, greenish black mud, 2 cm pebble at 45 cm, greenish black layer (organic ?) at 49 to 51 cm
60		SAND: very fine sand laminated with mud, diffuse upper and lower contacts
		LAMINATED MUD: with sand laminae, dark greenish gray
70		SAND: fines upward from medium to fine sand, scoured base, greenish black
		SAND: very fine sand, sharp lower contact
		LAMINATED MUD: sand laminae increase at top of unit, mud is dark greenish gray, sand is darker
80		HOMOGENEOUS MUD: greenish black, (organic ?)
		LAMINATED MUD: dark greenish gray
90		LAMINATED MUD: 2 mm thick sand laminae at 87 and 89 cm, dark greenish gray

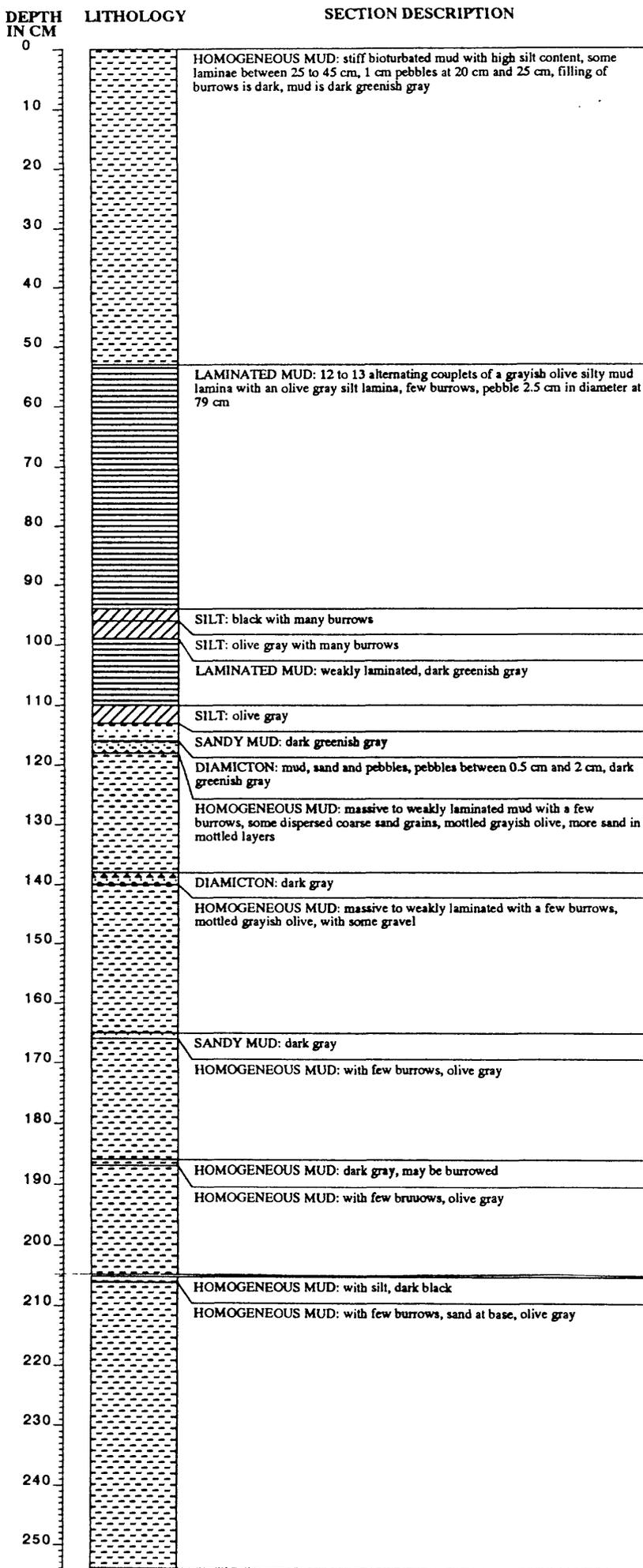
AH93 GB GC-11

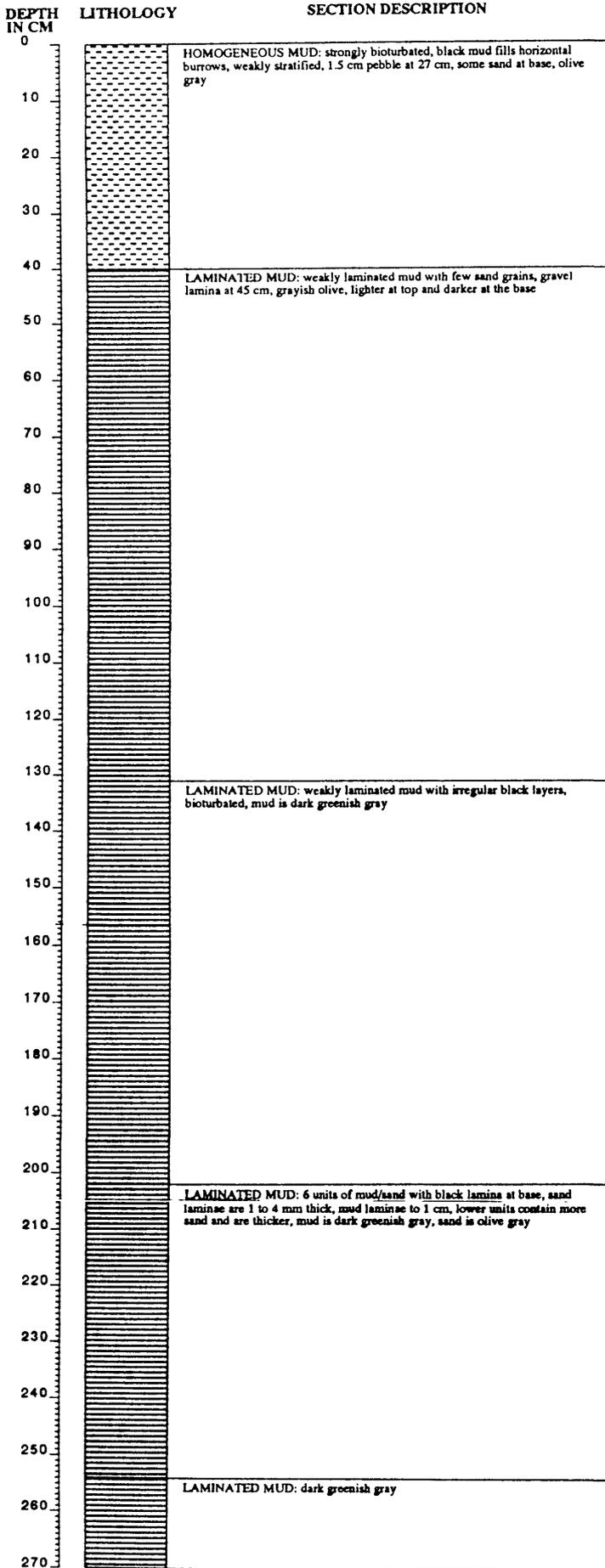
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		HOMOGENEOUS MUD: scattered sand grains, grayish olive
10		LAMINATED MUD: faint laminae, dark greenish gray
20		HOMOGENEOUS MUD: contains horizontal burrows, mud clasts, organic rich, olive gray
30		LAMINATED MUD: olive gray
40		HOMOGENEOUS MUD: very light grayish olive mud with a moderate olive brown patch
50		LAMINATED MUD: with scattered sand grains, 2.5 cm pebble with sand grains around it at 47 cm, olive gray
60		HOMOGENEOUS MUD: black (organic ?)
70		LAMINATED MUD: olive gray
80		LAMINATED MUD: olive black mud above grayish olive with olive gray beneath, black mud may be organic
90		LAMINATED MUD: olive gray, 4.5 cm pebble at 80 cm with gravel around it
100		SAND: medium sand
110		LAMINATED MUD: with several thin sand stringers greater than 1 mm, color is various shades of grayish olive
120		LAMINATED MUD: interlaminated fine sand and mud, sand laminae are up to 1 mm, mud laminae up to 2 mm, grayish olive
130		HOMOGENEOUS MUD: black (organic ?)
		SANDY MUD: grayish olive

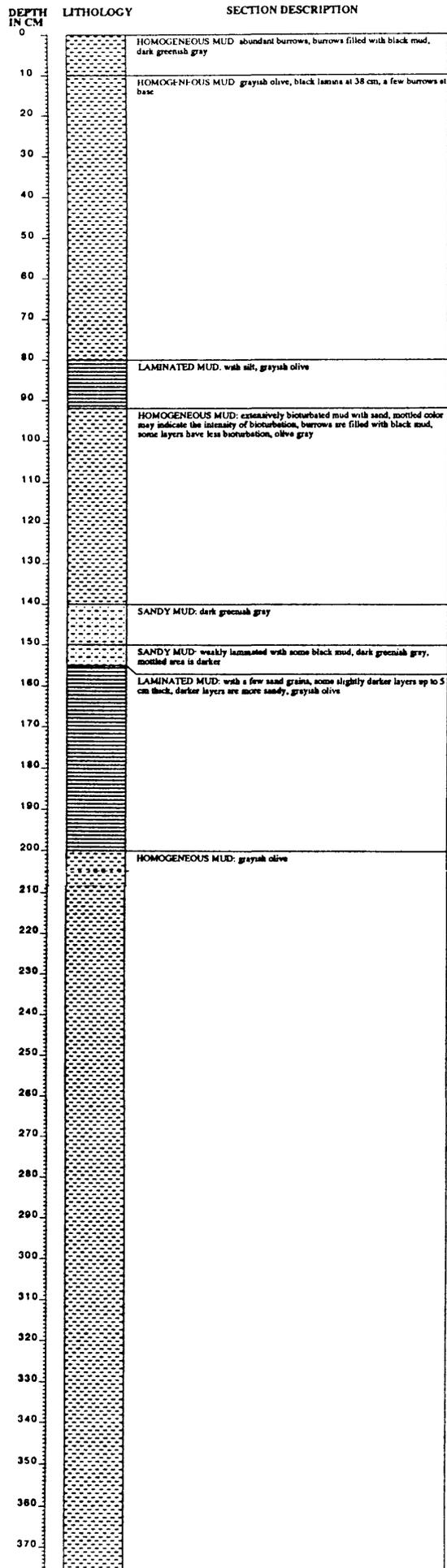


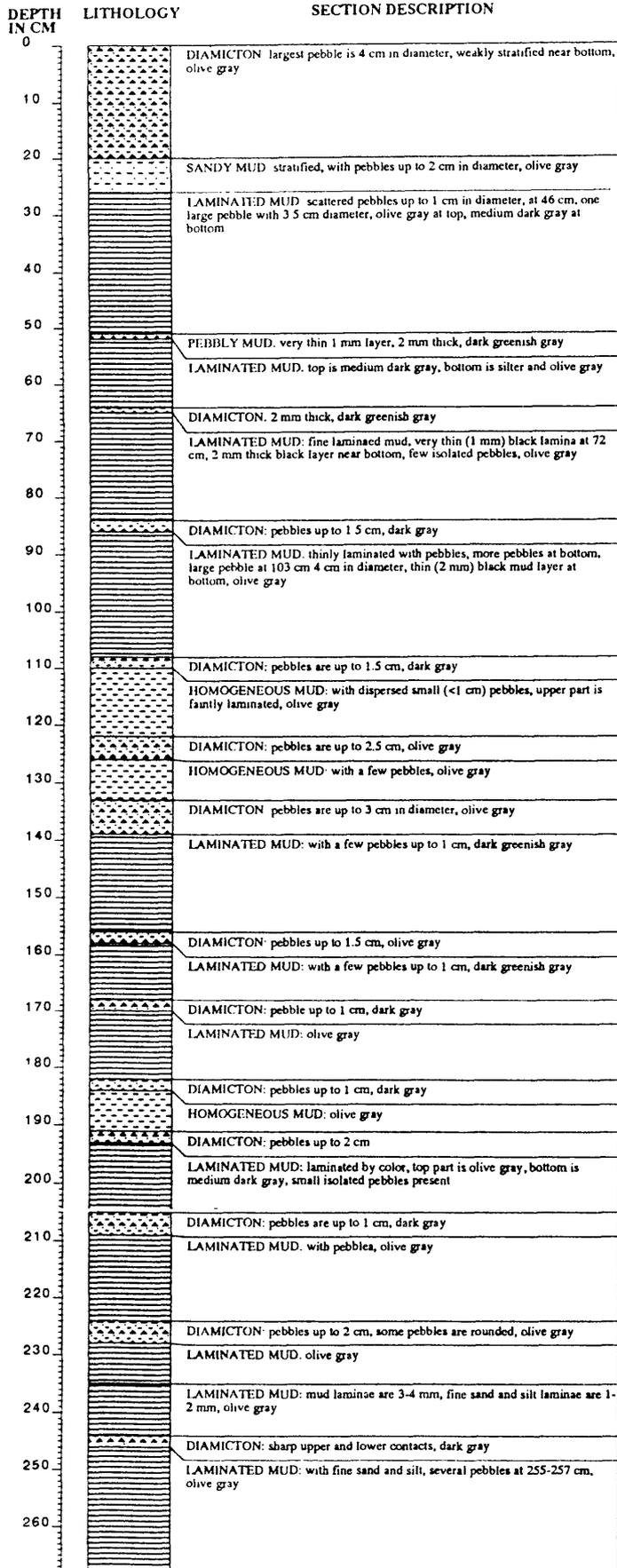


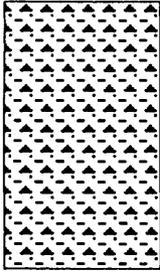


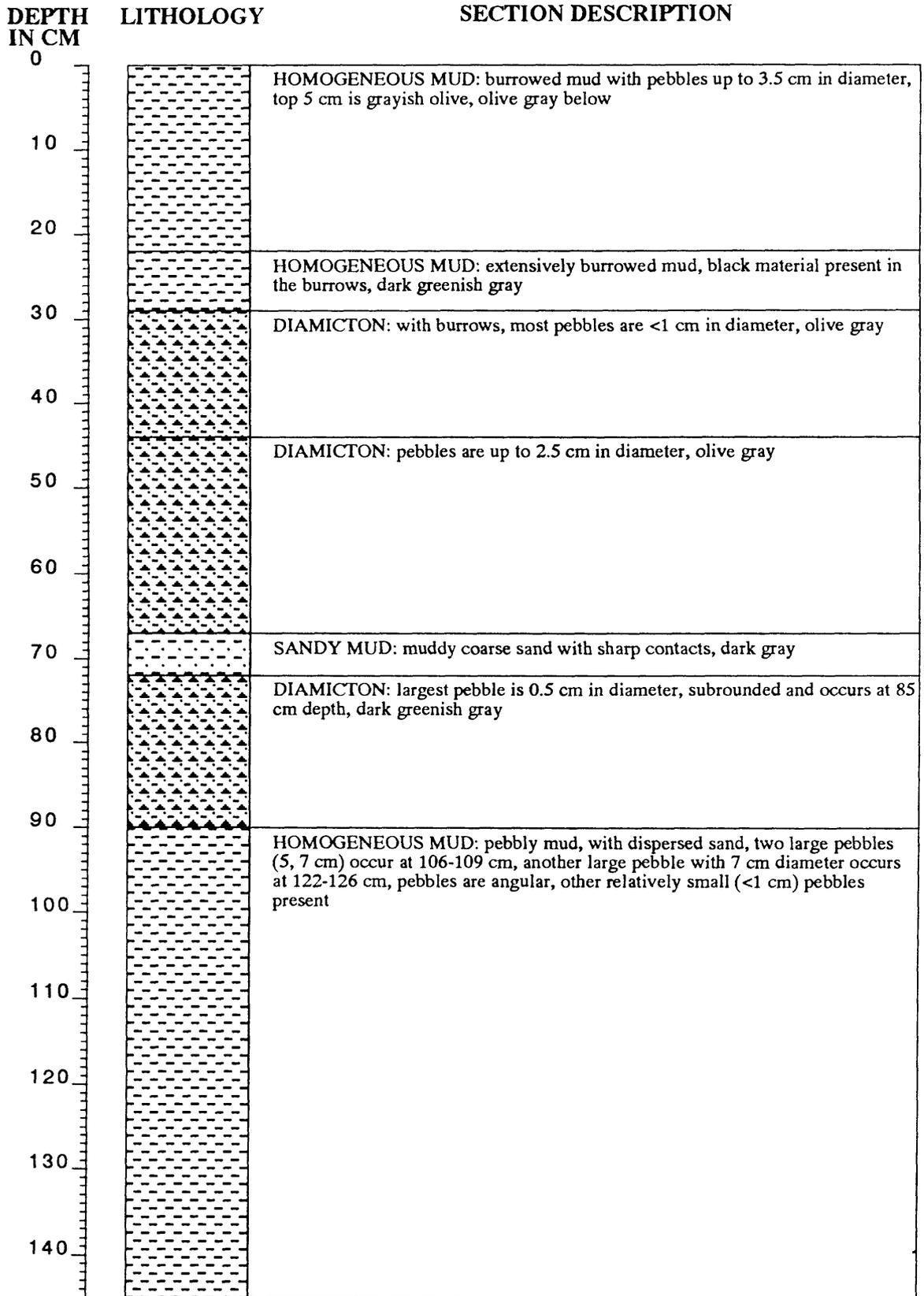


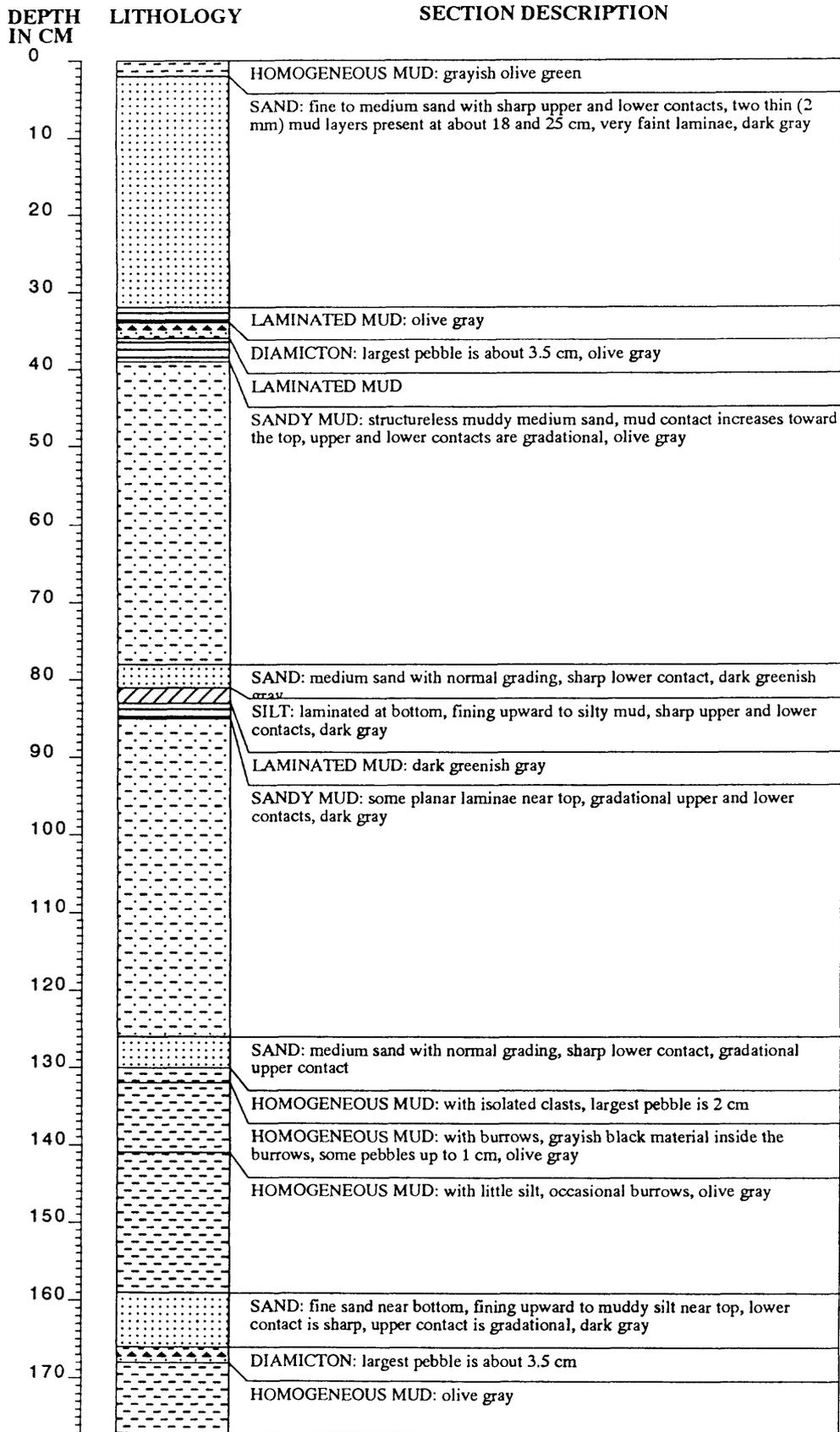


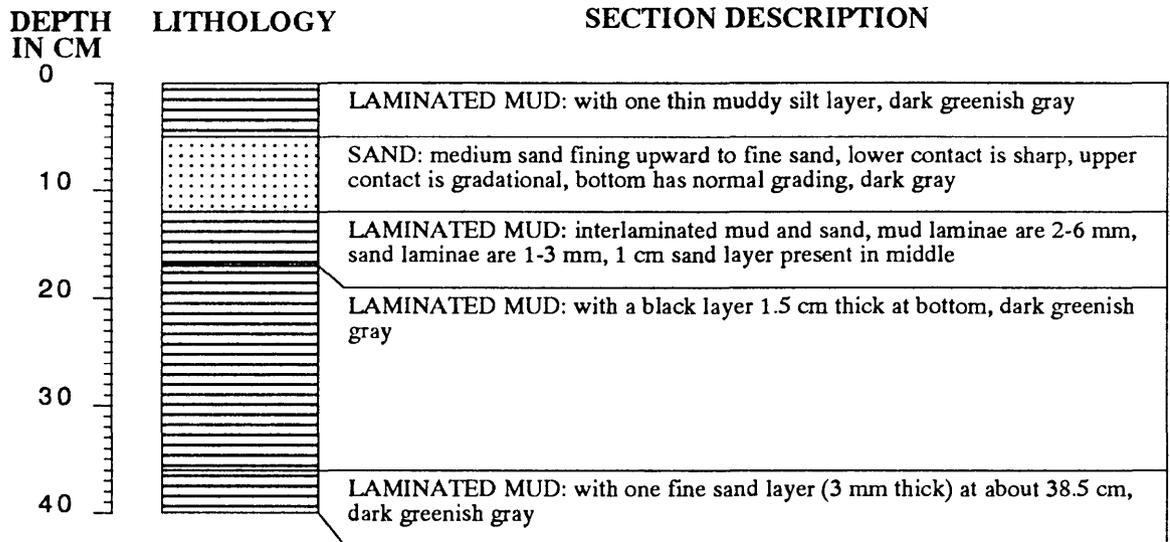




DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0 10 20		DIAMICTON: largest pebble is 4 cm, pebbles are subrounded to angular, worm tubes present near top, dark greenish gray

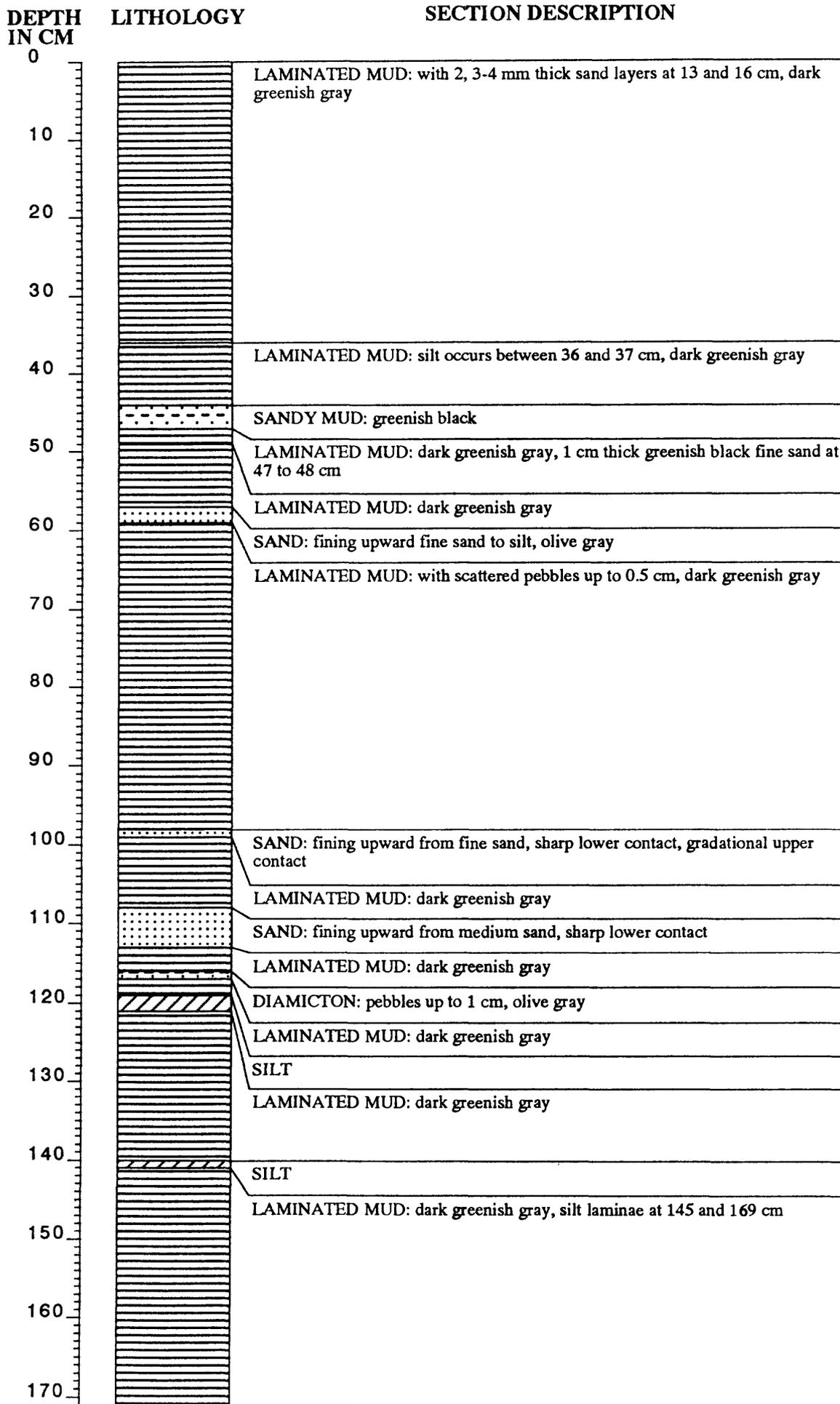


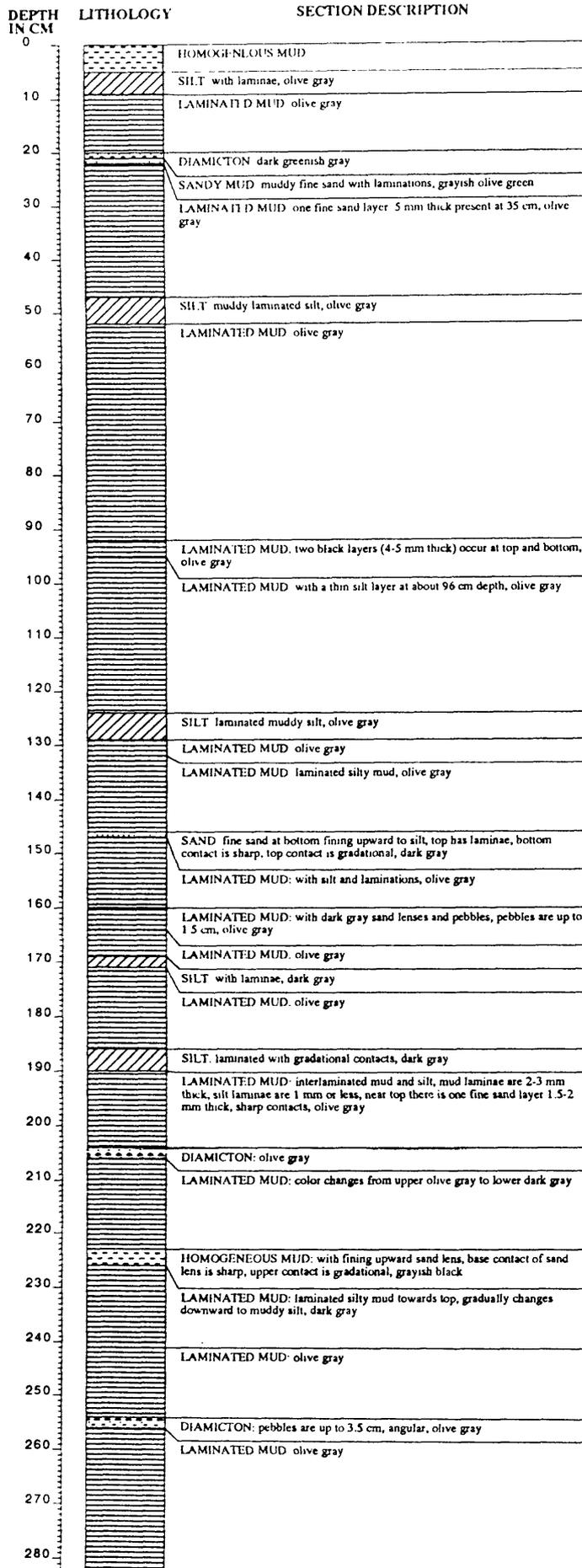


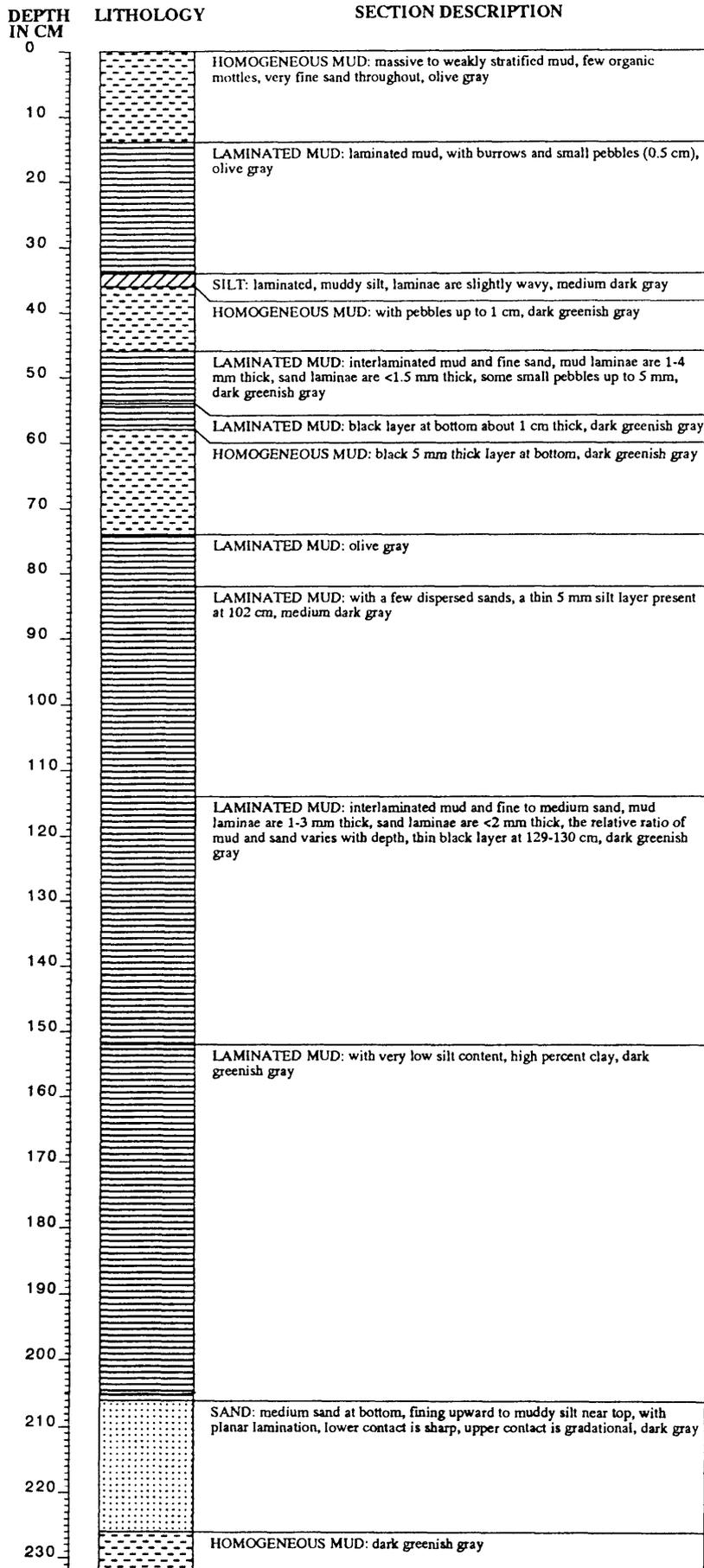


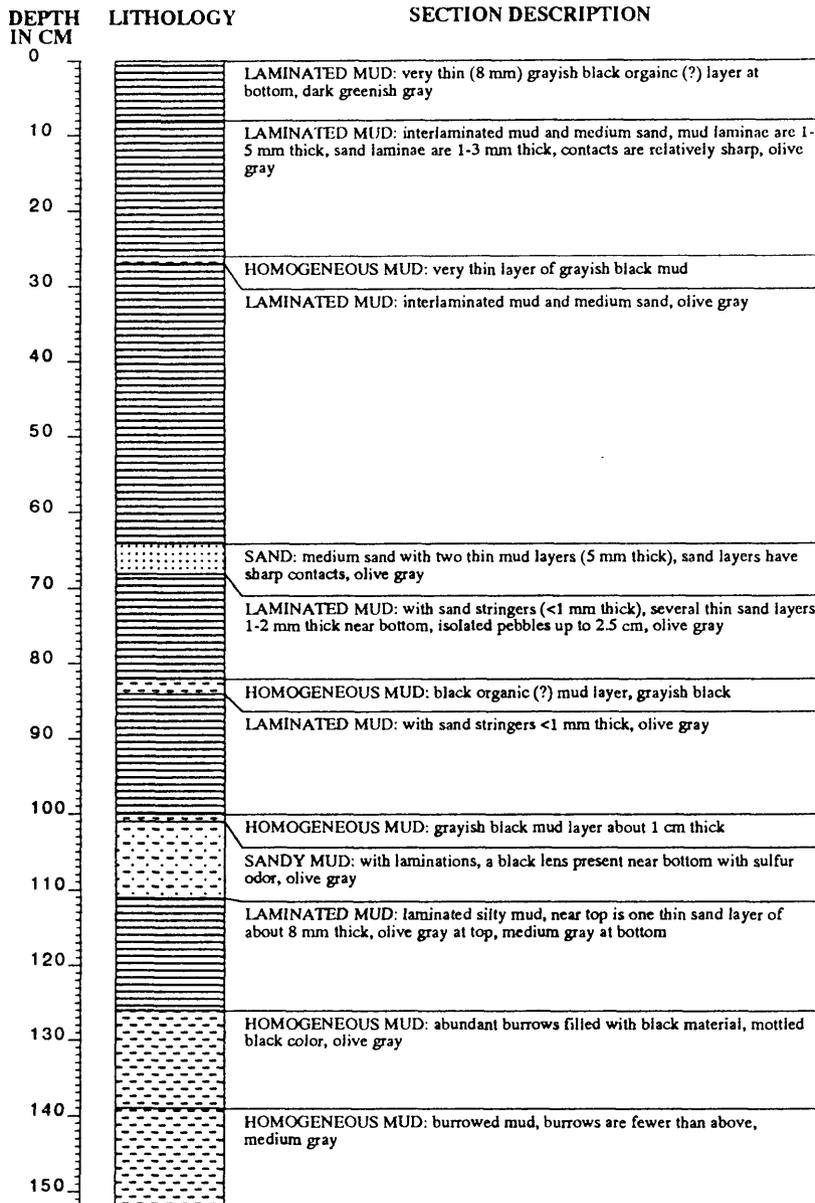
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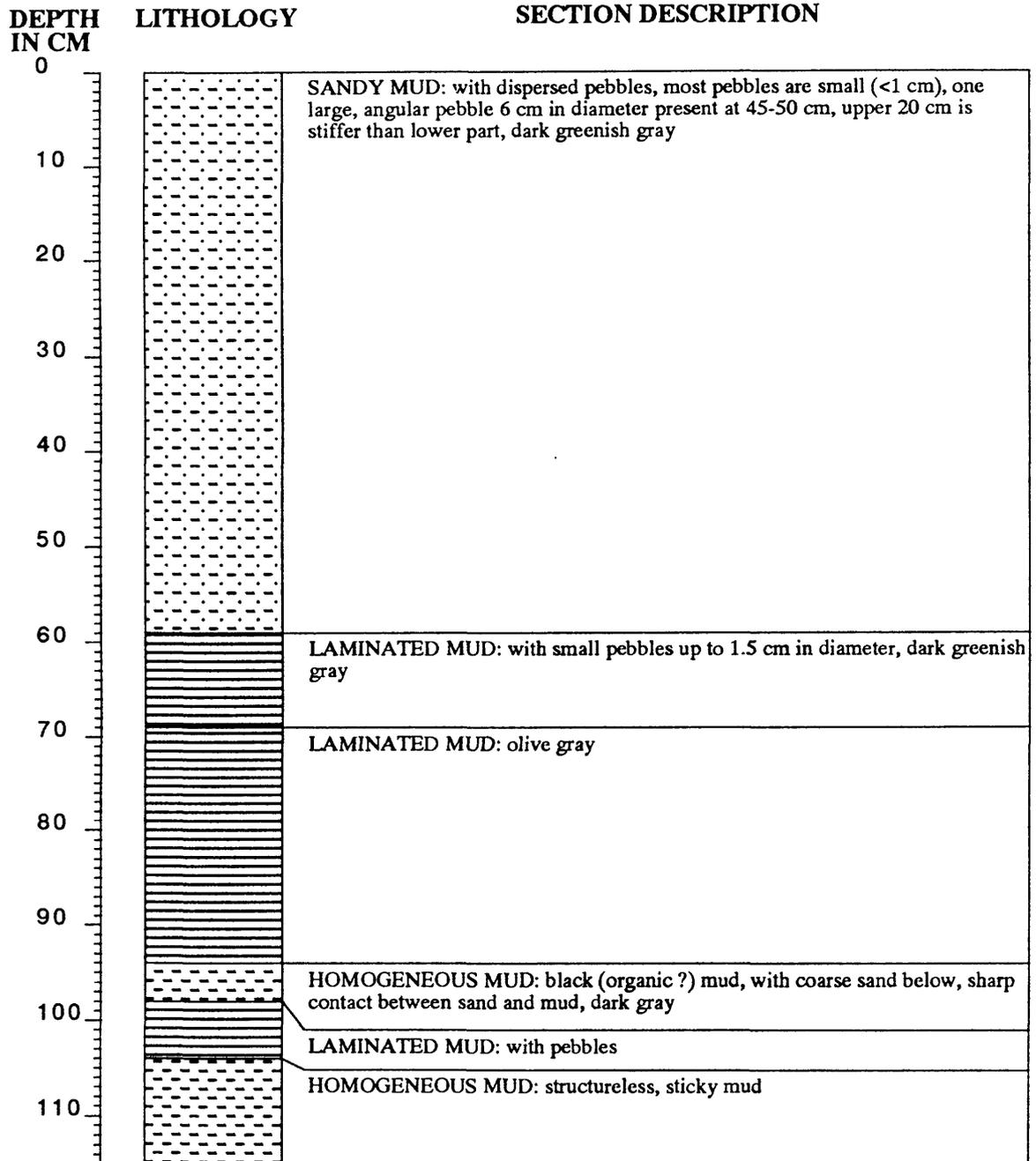
DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		PEBBLY MUD: mud with angular pebbles up to 4 cm, becomes sandy near bottom, olive gray
10		PEBBLY MUD: pebbles are up to 3 cm in diameter, some worm tubes present between 15 and 20 cm, olive gray
50		SANDY MUD: with some small pebbles (<1 cm), olive gray



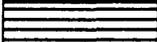






AH93 GB GC-33

AH93 GB GC-34

DEPTH IN CM	LITHOLOGY	SECTION DESCRIPTION
0		LAMINATED MUD: weakly laminated with burrows, small pebbles up to 1 cm, dark greenish gray
10		LAMINATED MUD: laminae are up to 1 mm thick, grayish olive
20		LAMINATED MUD: interlaminated mud with olive gray and dark greenish gray colors, pebbles up to 1 cm, 5 cm coarse sand layer at 46 cm, black layer 6 mm thick at 51 cm
30		LAMINATED MUD: dark greenish gray
40		LAMINATED MUD: dark greenish gray
50		LAMINATED MUD: dark greenish gray

APPENDIX 2: PHYSICAL PROPERTIES MEASUREMENTS AND GRAPE LOGS

Physical Properties Measurements of Sediment Sampled from Western Prince William Sound, Yakutat Bay, and Glacier Bay, Southeastern Gulf of Alaska.

R.E. Kayen

Cores sampled during AH93 were logged for their physical properties on a multi-sensor whole core sediment logging device. The physical property logging device was built in Great Britain by Schultheiss Geotek, Ltd. This device is controlled by a personal-computer driven software system, developed at the U.S. Geological Survey, Branch of Pacific Marine Geology (USGS-BPMG), for data acquisition and instrument I/O manipulation. The system logs sediment bulk density, compression wave sound speed, and magnetic susceptibility of unsplit whole sediment cores. These properties are used to develop physical property profiles for the cores.

Whole-section sealed sediment cores are placed horizontally upon a transport sled. This sled is transported through a frame supporting sensors in incremental fashion by a computer-controlled stepper motor. Sensor readings of cores were logged in this manner at 1 centimeter increments, within the first day after the cores were collected. The transport sled is capable of carrying core sections up to 1.50 m long.

Sediment bulk density (ρ_b) is calculated from the gamma ray attenuation characteristics of the cores according to Lambert's law. The configuration of our device allows for a core to pass between a vessel containing the radio-isotope Cesium-137 and a scintillation counter. For a defined time period, the number of gamma decays emitted from the Cesium-vessel that pass through the core and are received at the scintillation detector are counted. If gamma rays are transmitted to the scintillation counter through air, we refer to the number of scintillation counts as an unattenuated measure, I_0 . For the case where a material of some thickness, d , lies between the Cesium source and sensor, the attenuated gamma ray counts, I , can be related to the unattenuated number of gamma decays, I_0 , the material thickness, d , the material bulk density, ρ_b , and the Compton scattering coefficient, μ , by Lambert's Law:

$$I = I_0 \exp\{-\mu\rho_b d\}$$

The bulk density of the material can be determined as follows:

$$\rho_b = \frac{1}{\mu d} \ln \left(\frac{I_0}{I} \right)$$

For whole sediment cores, we need to account for the influence of the core liner to get an accurate estimation of the sediment density. As such, we determine separate Compton scattering coefficients, material thicknesses, and bulk densities, for the sediment and core liner. Doing so, we can determine the sediment bulk density from the following equation:

$$\rho_b = \frac{\left(\ln \left(\frac{I_0}{I} \right) - 2L \rho_{\text{liner}} \mu_{\text{liner}} \right)}{\mu_{\text{sed}}(D - 2L)}$$

where D is the whole core outer diameter (Sediment diameter plus two liner wall thicknesses), L is the liner thickness, ρ_{liner} is the liner density, μ_{liner} is the liner Compton scattering coefficient, and μ_{sed} is the sediment Compton scattering coefficient.

The compression wave velocity, V_p , of sediment is calculated from the measured core diameter and wave travel time, correcting for the liner thickness, electronic signal delays, and core liner travel time. The velocity is calculated as:

$$V_p = \frac{D-2L}{T - 2T_{\text{liner}} - T_{\text{delay}}}$$

Parameters for the velocity calculation are the total travel time, T , the liner travel time, T_{liner} , and the electronic signal delay within the transducer, T_{delay} , all measured in μsec .

The magnetic susceptibility of sediment is measured directly through a transducer hoop. No liner corrections are required on cores collected during cruise AH93, as the liners are composed of non-magnetically susceptible polymers.

System Calibration

Calibration of the logger system with standards is required in order to present physically meaningful values of sediment properties. For each of the properties we measure, we have developed a set of standards to correctly attune the transducer and computer system to output calibrated physical property data.

Density & Velocity Calibration:

Density and compression-wave velocity measurements of whole core sediment are calibrated to the known standards of water and aluminum. These two standards serve as end-members which fully bracket seafloor sediment density, with water serving as the lower-bound and aluminum serving as the upper-bound. These standards also serve to nearly bracket sediment compression-wave velocity (near surface fine-grained deposits of low density may have velocities below that of water.). The added advantage of using these materials is that their respective Compton scattering coefficients are similar to those of sediment which is composed of a liquid-phase of water and a solid-phase, typically, of alumina-silicate minerals. To account for the influence of the liner, the Compton

scattering coefficient for empty core liner was determined prior to cruise AH93, by measuring the attenuation of gamma rays transmitted through the liner relative to the unattenuated air count, measuring the liner thickness and density, and using Lambert's law, above.

A water-aluminum standard was prepared by inserting a solid-cylinder of 6250-Aluminum into an unsplit section of core liner identical to the liner used for sediment sampling. The length of milled aluminum fills one-half the total length of the 'calibration standard'-core liner and distilled water fills the remaining portion. During density calibration, the numbers of scintillation's-per-second were logged during transmission of gamma rays through the liner and water, the liner and aluminum, and through air. Finally, an empirical Compton scattering coefficient was determined for the water, and aluminum which gave water densities of 1.00 g/cc and aluminum densities of 2.70 g/cc. Calibration studies, performed in the geotechnical engineering laboratory of the USGS-BPMG, indicate that the standard deviation for density measurements is on the order of 0.6-1.0% of the standard measured value (1.00 and 2.70 g/cc).

Calibration standards were run repeatedly during the logging of the cores, onboard the Alpha Helix. That is, in order to calibrate the sediment-core profiles for density, we took attenuated gamma-ray measurements from our water and aluminum filled liner standard, and unattenuated measurements through air, after every core section was logged on our device. For each sediment-core file, we applied the same Compton scattering parameters that corrected the corresponding calibration-file standards, water and aluminum, to their known values to the gamma decay-counts measured through the sediment collected from the fjords of the Southeastern Gulf of Alaska.

To calibrate compression wave velocity measurements in our water-filled standard, we needed to measure the temperature of the water. We then corrected the raw calculated velocity to the velocity at 23°C at standard pressure using known correction factors (U.S. Naval Oceanographic Office, 1962). The known velocity of distilled water at standard pressure and temperature is 1.4917 km/sec. An empirically determined travel time delay ($2T_{\text{liner}} - T_{\text{delay}}$) was determined which corrects the measured p-wave velocity to the known standard velocity. This empirical travel-time delay, from the calibration files, was then applied to the corresponding sediment-core file. Prior studies indicate that the standard deviation of our sediment velocity measurements is on the order of 0.16% of the measured value.

Magnetic Susceptibility Calibration:

Calibration of magnetic susceptibility measurements is done by suspending a standard of known susceptibility within the magnetic susceptibility hoop transducer and determining the correct linear proportionality constant which scales the measured value of the standard to the known value. The known standard that we used to calibrate the system was supplied by the manufacturer of the logger system. This correction factor was then used to adjust the measured magnetic susceptibility of sediment to a calibrated value. Unlike the discrete measurements made for density and velocity at each one centimeter interval in a core, the magnetic susceptibility hoop senses a broader section of the core (several centimeters outside the hoop). This sensitivity has the effect of smoothing the magnetic susceptibility curves and reducing the logger resolution of fine scale variations of magnetically-influenced minerals (magnetite, etc.) in the sediment cores.

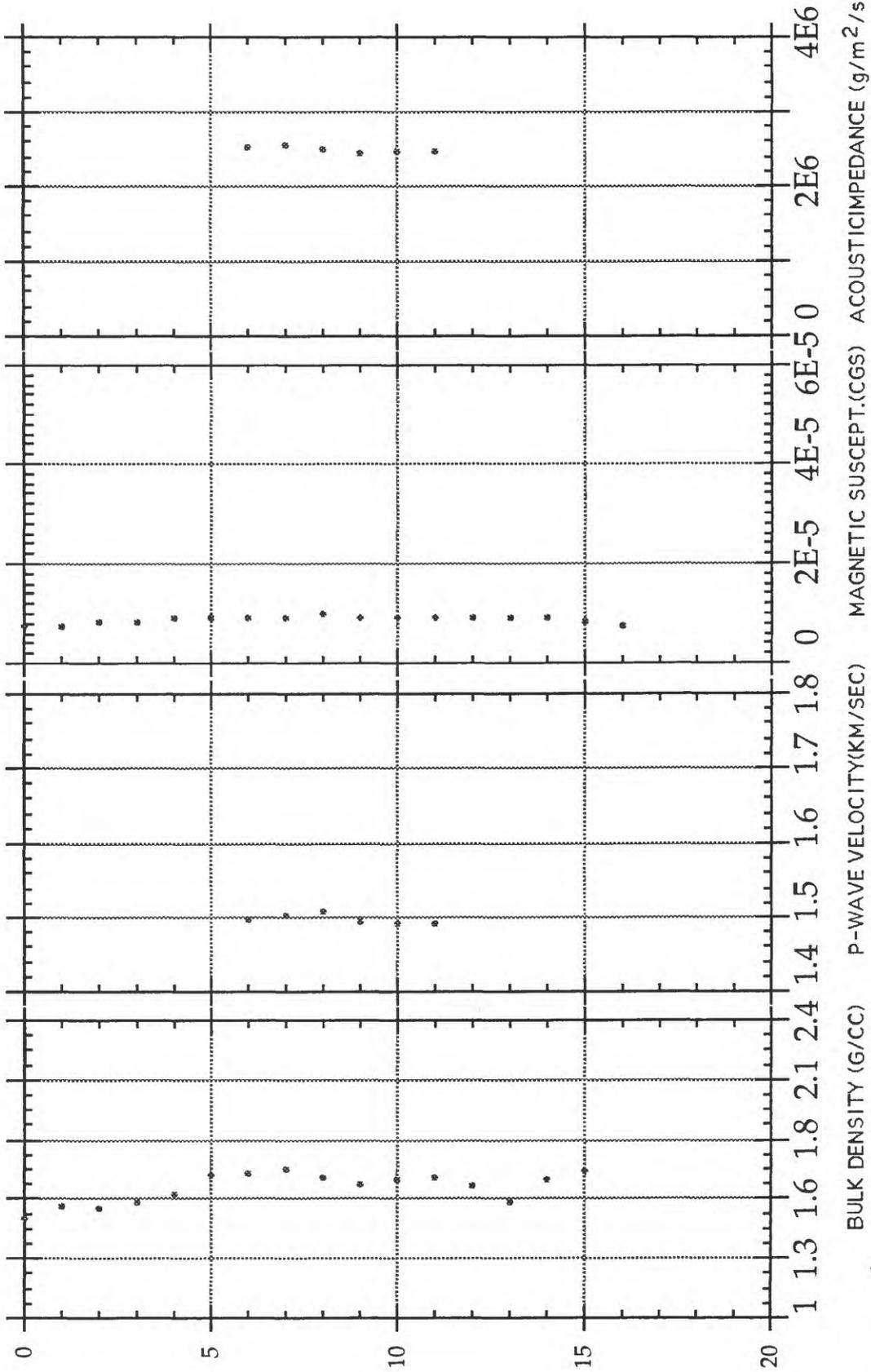
References:

CRC Handbook of Chemistry and Physics, 1969, R.C. Weast, ed. The Chemical Rubber Co., Cleveland, Ohio.

U.S. Naval Oceanographic Office, 1962, TABLES OF SOUND SPEED IN SEA WATER, Oceanographic Analysis Division, 47p. (SP-58).

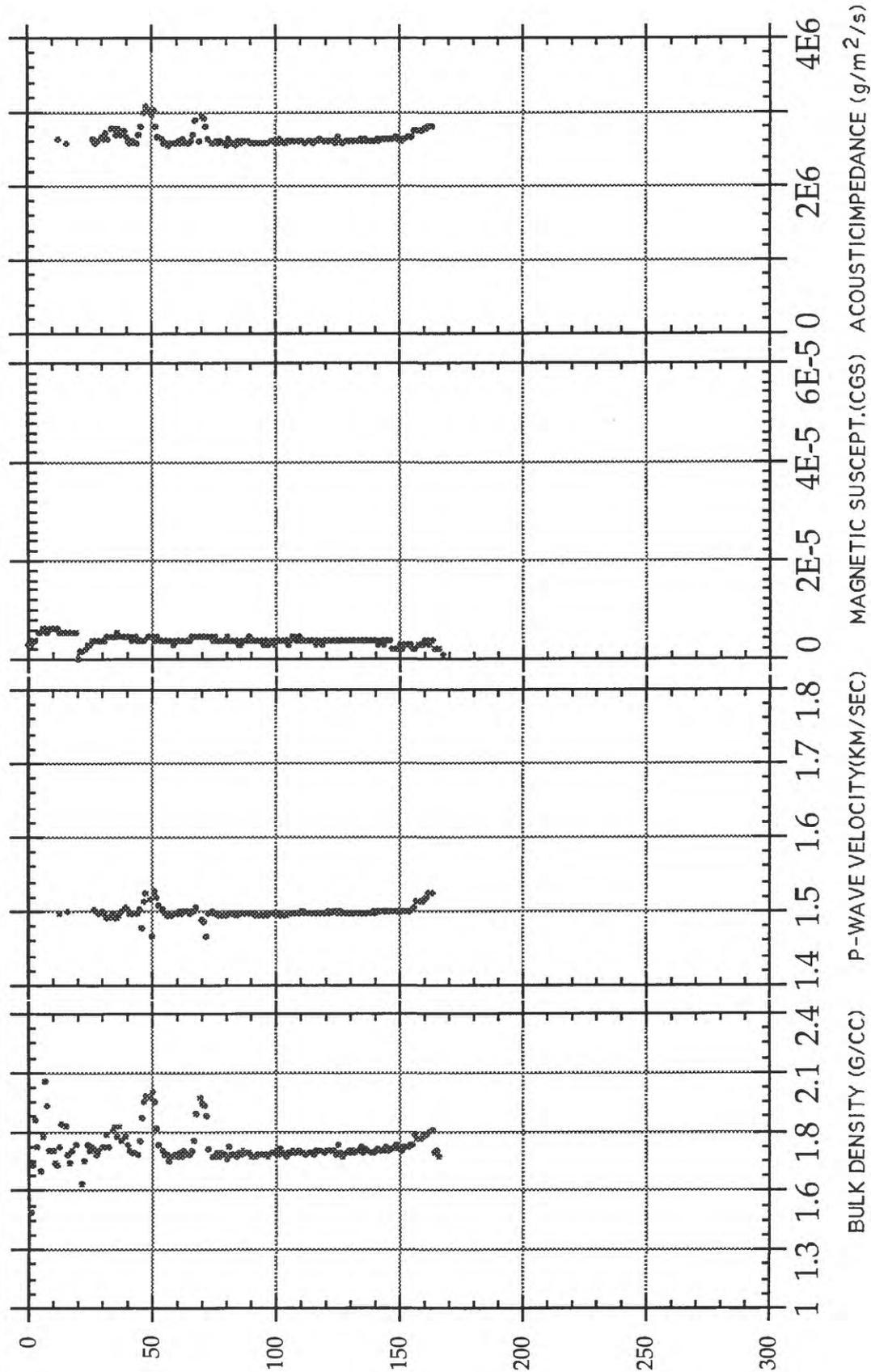
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AH93-PW-B1



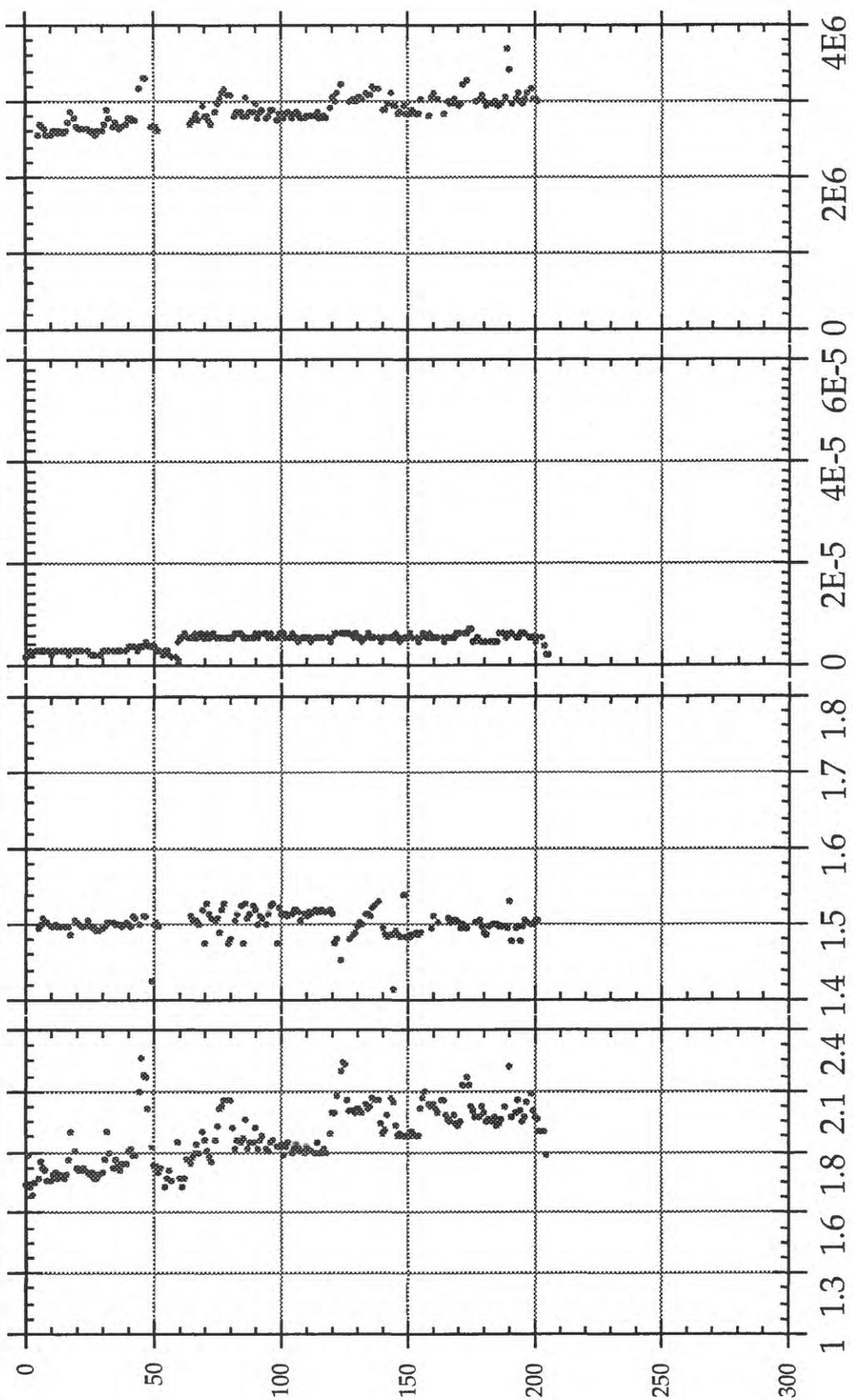
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AH93-PW-GC1



AH93-PRINCE WILLIAM SOUND

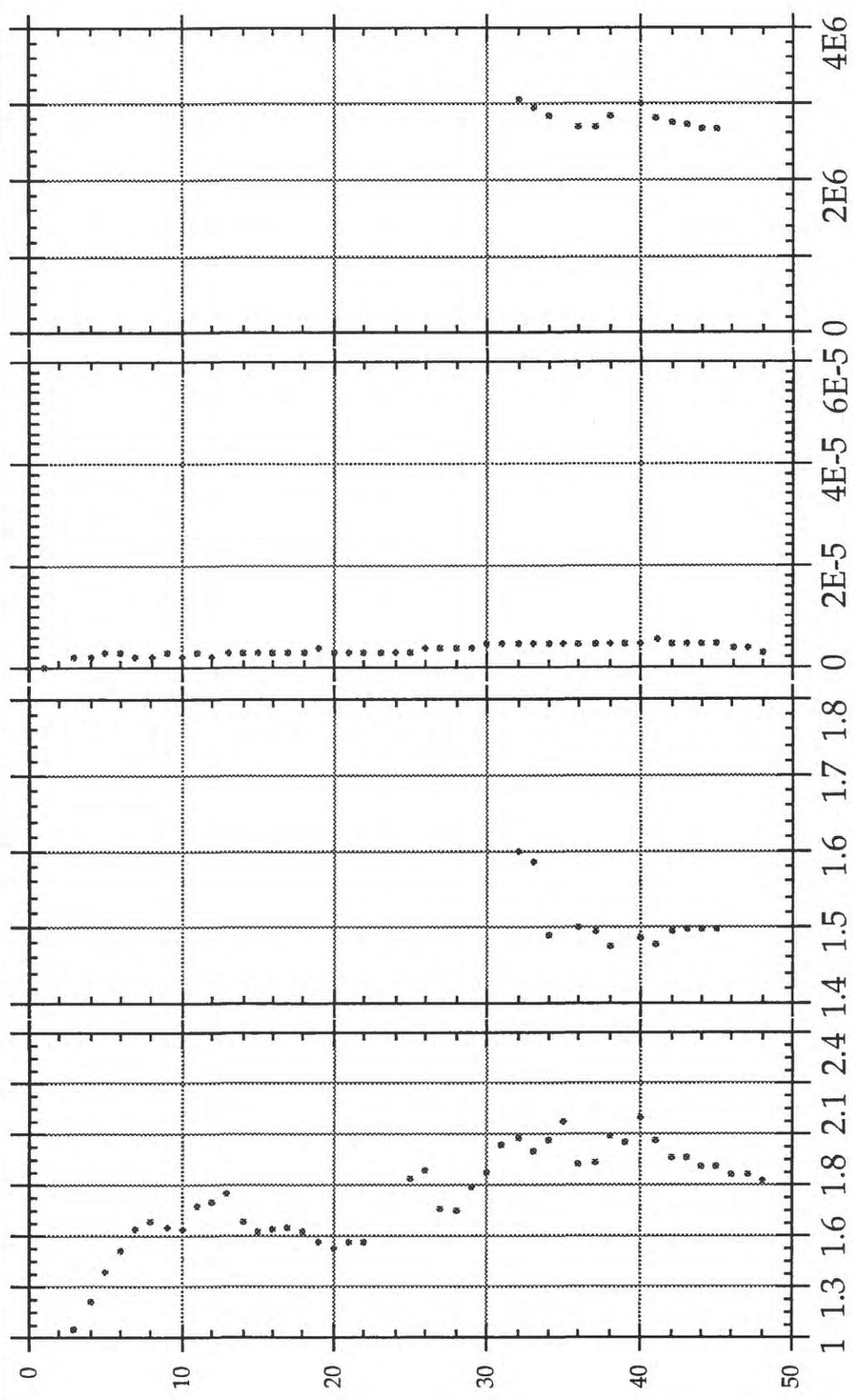
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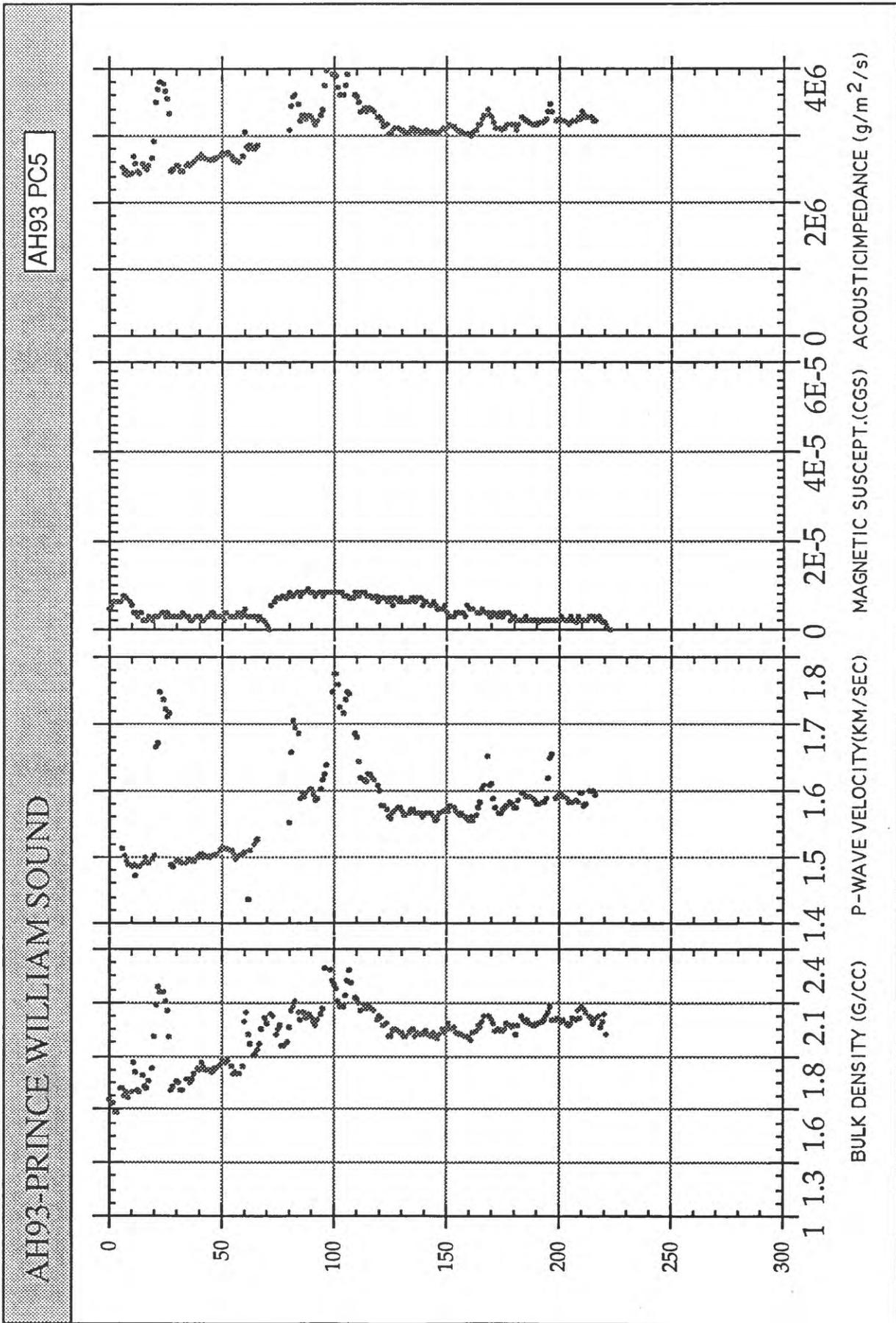
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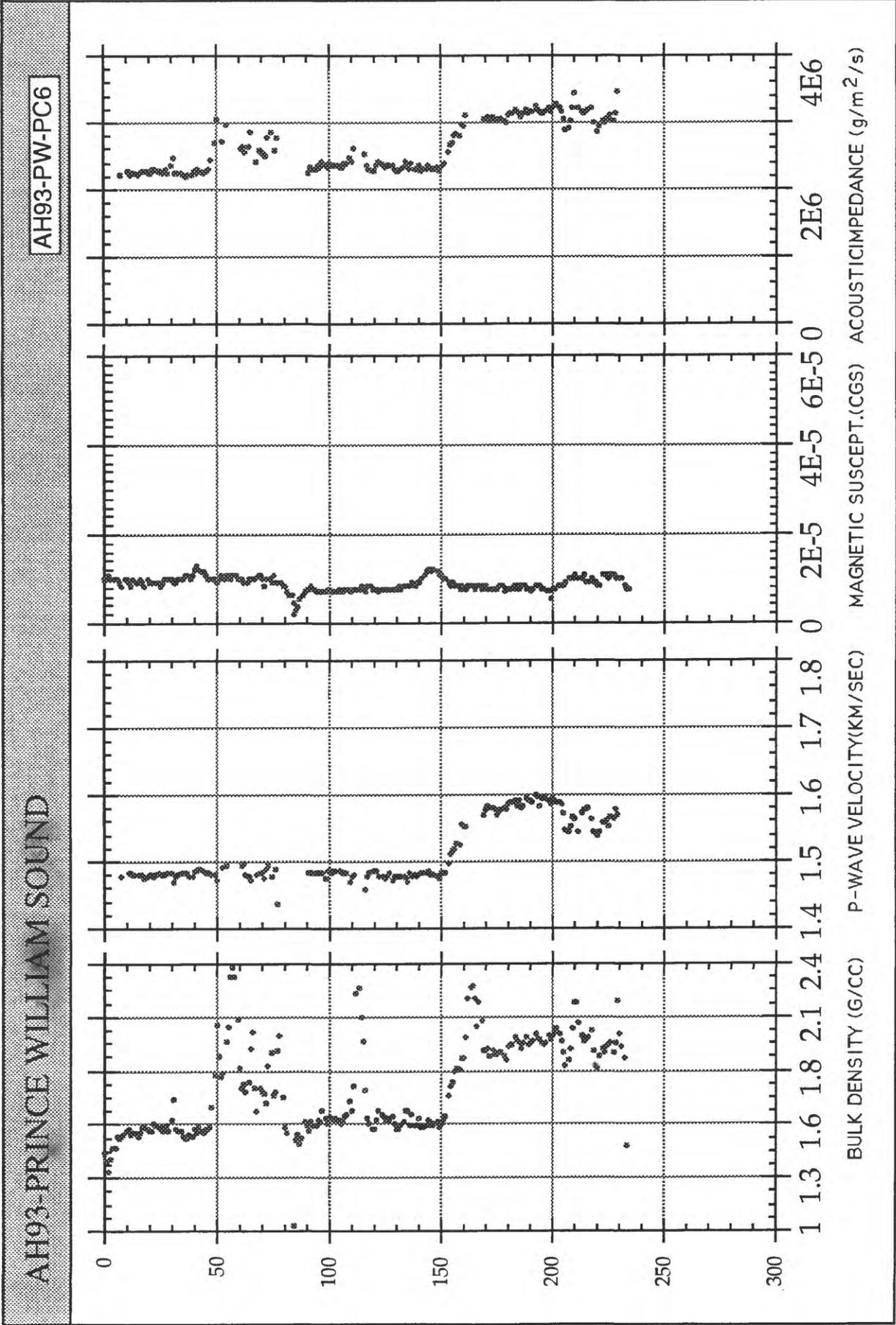
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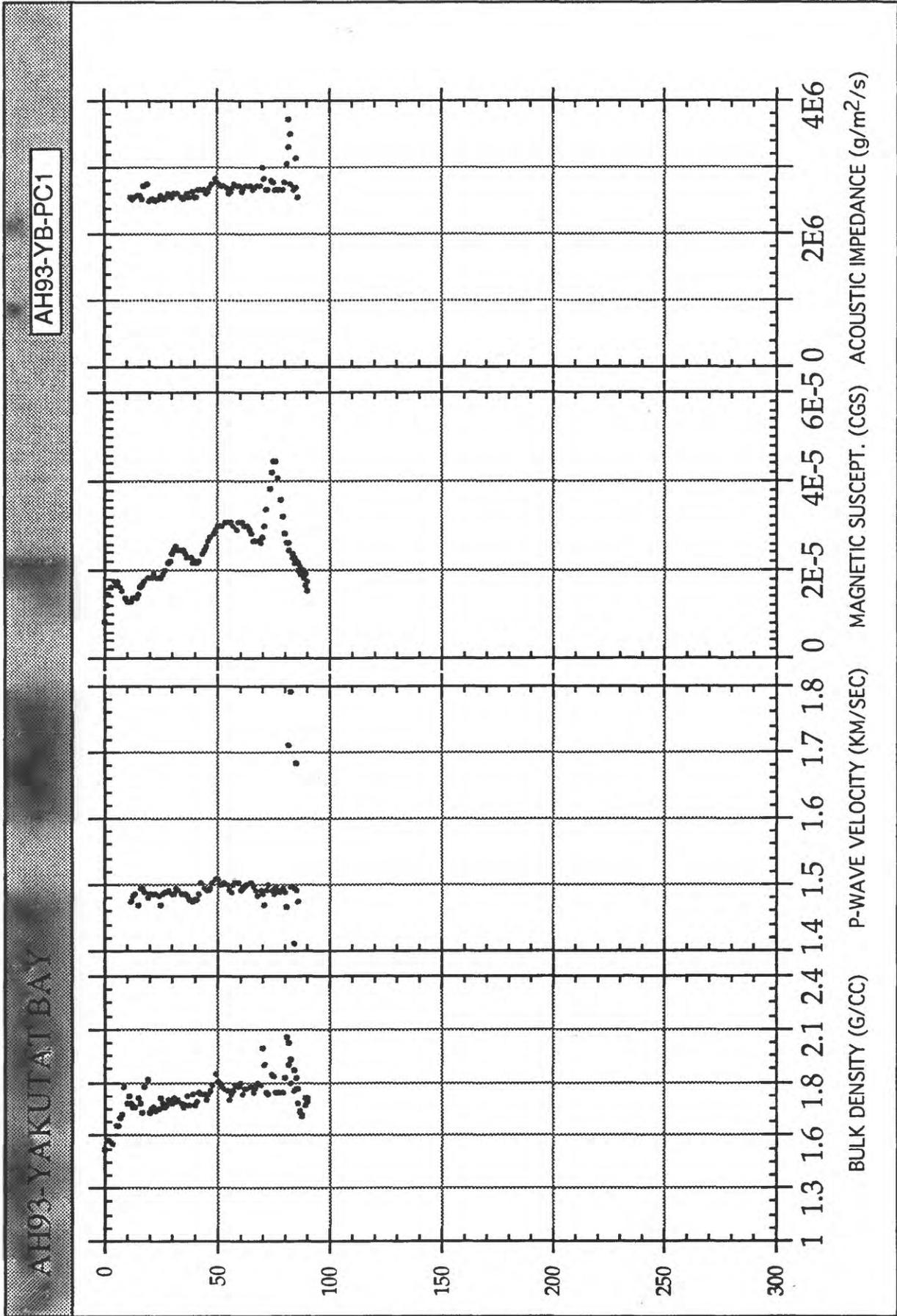
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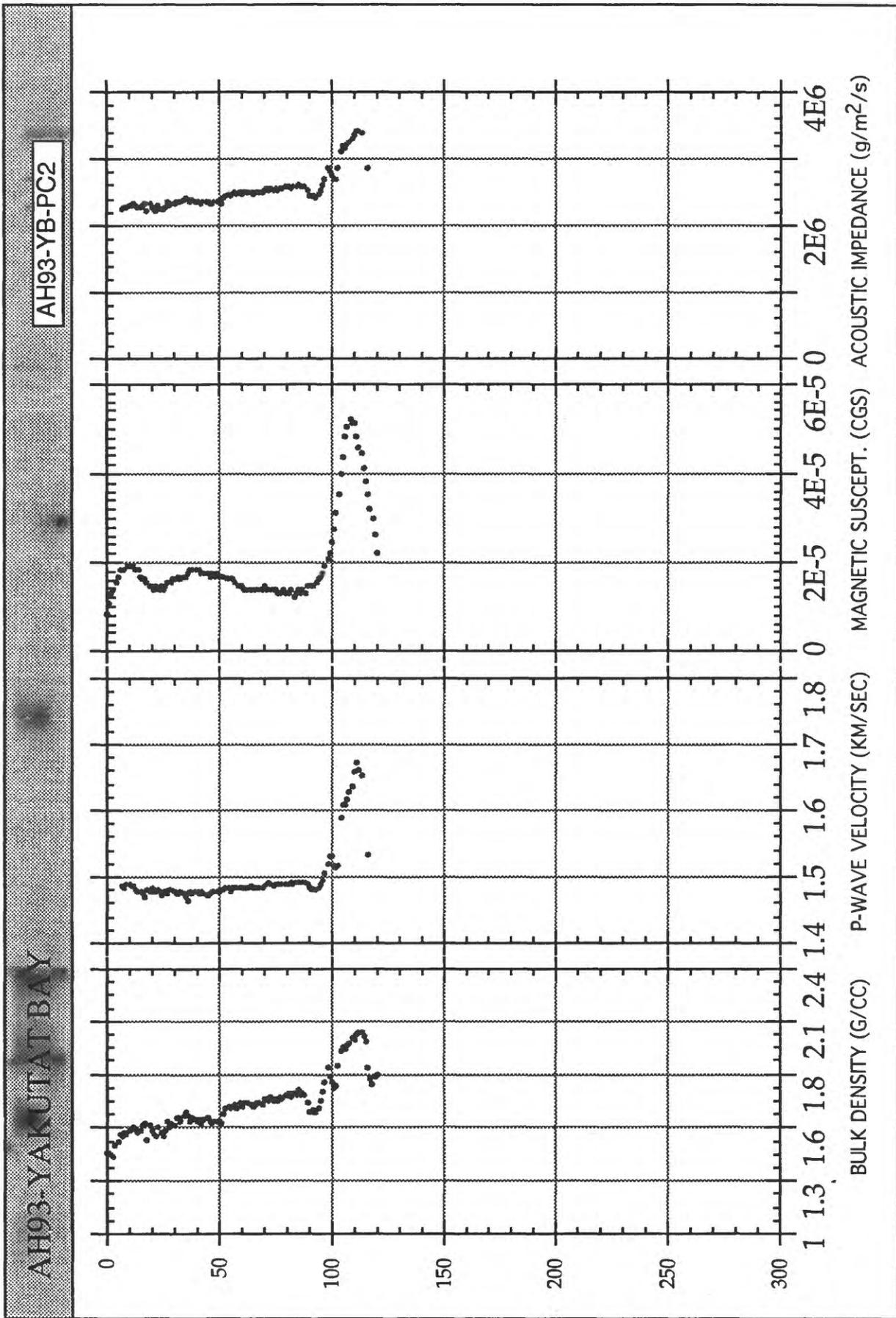


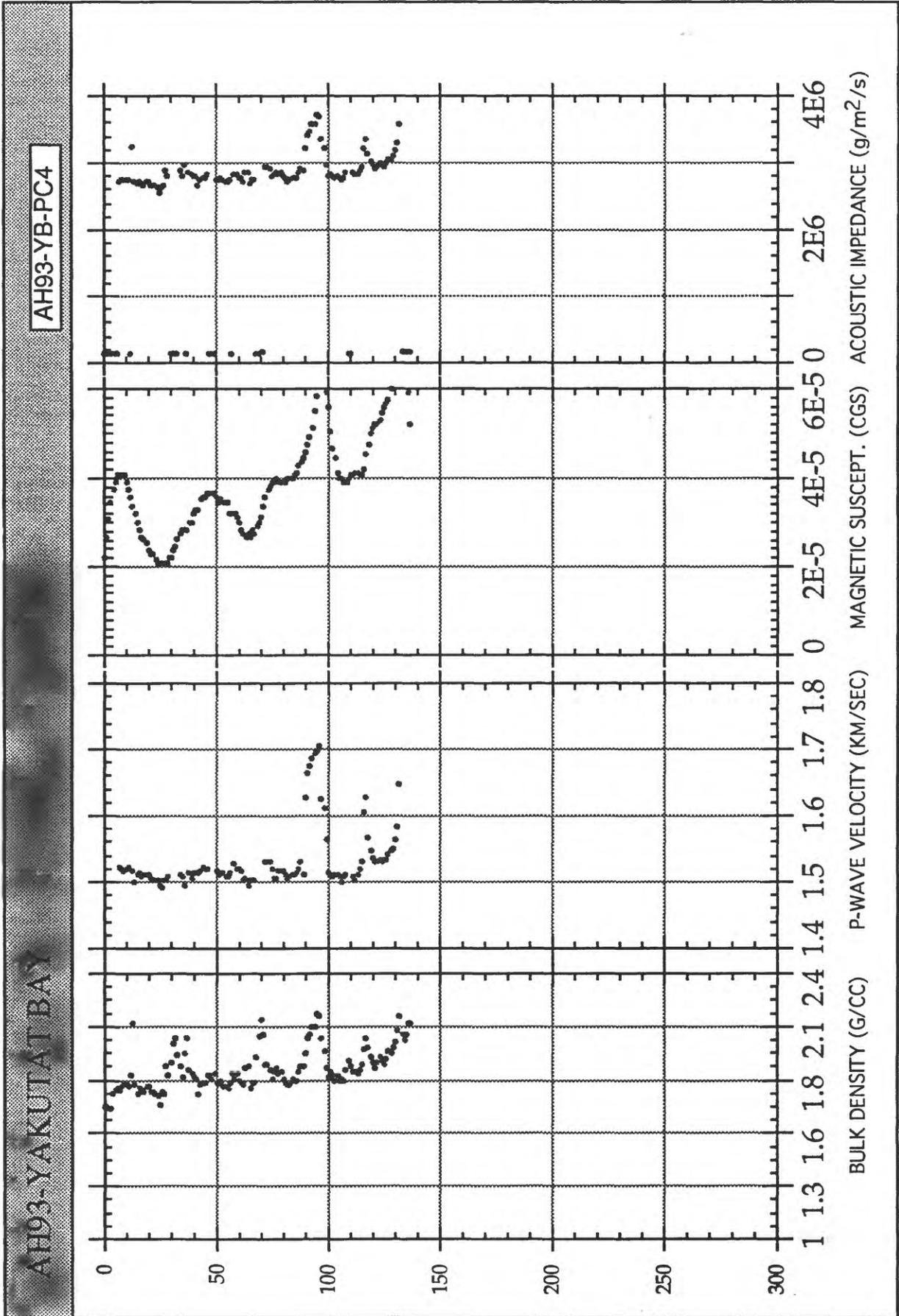
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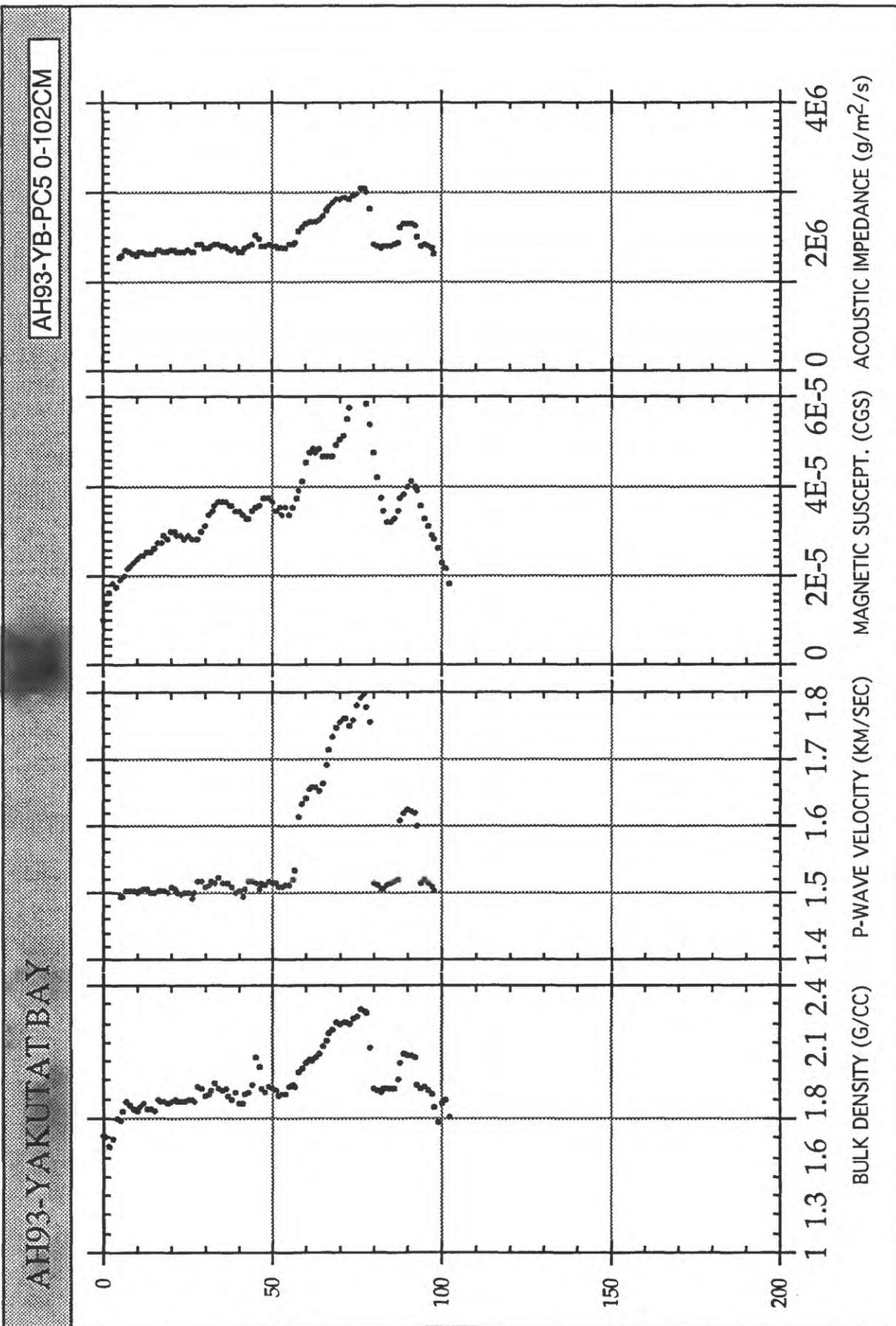


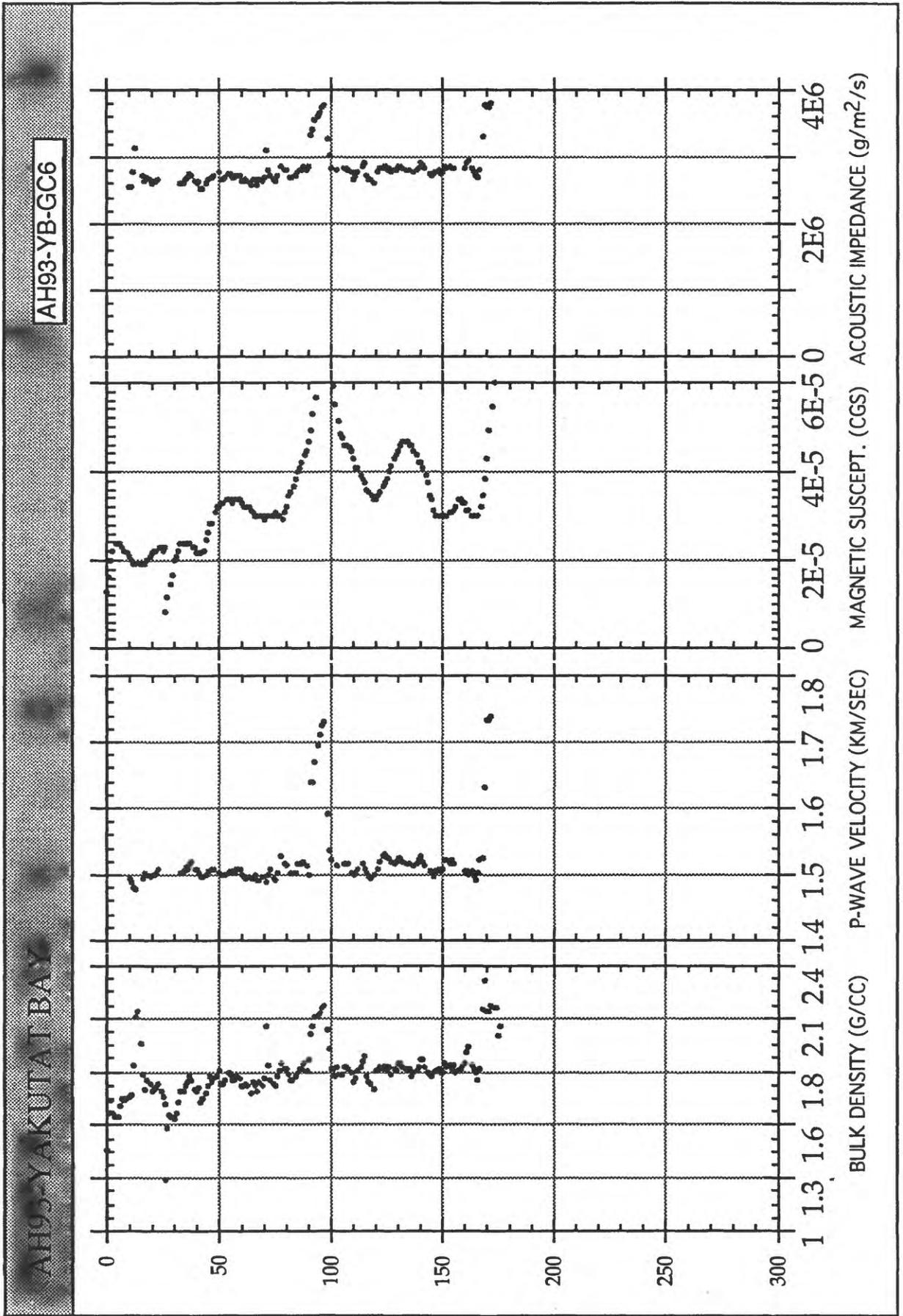






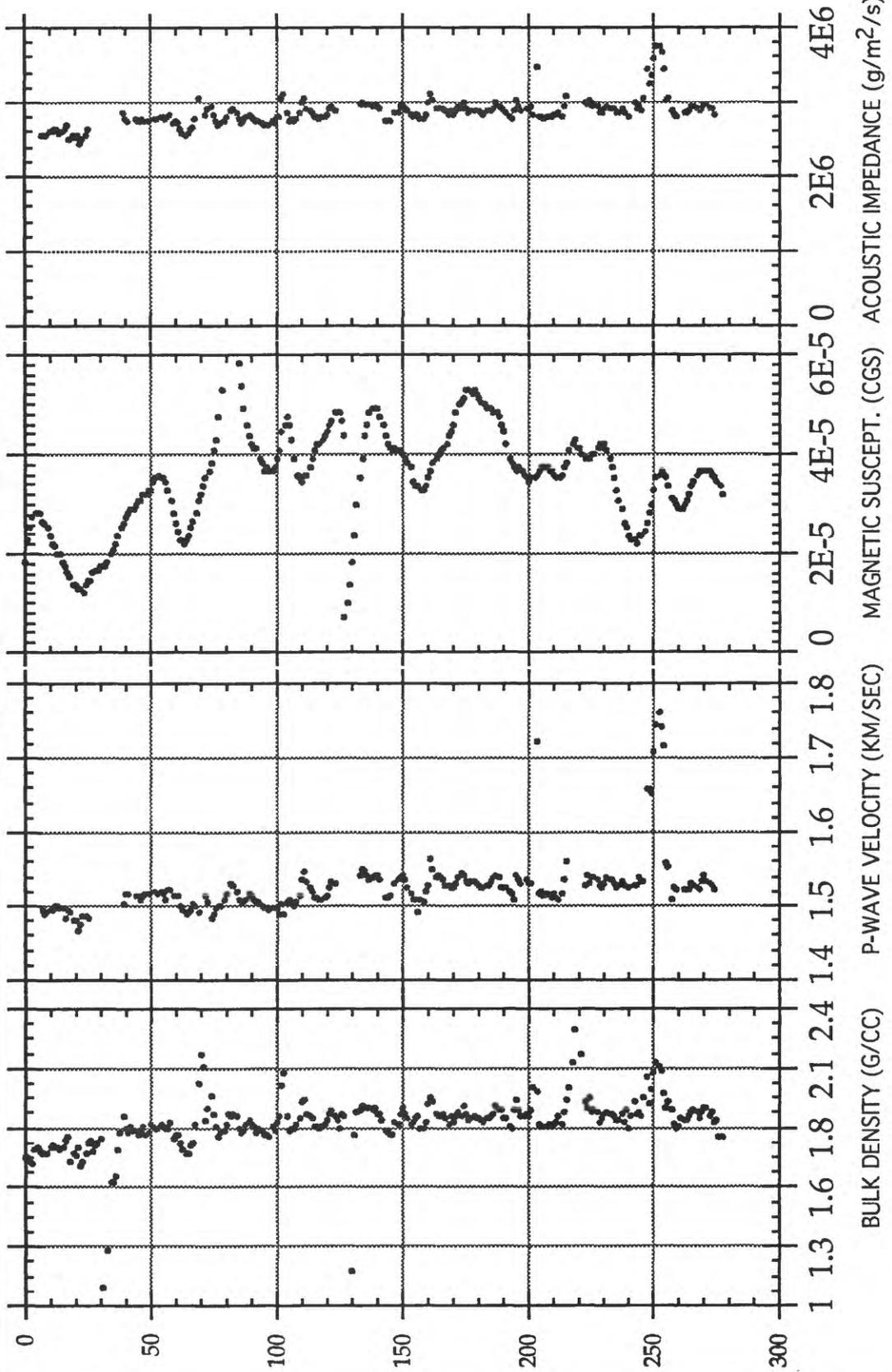


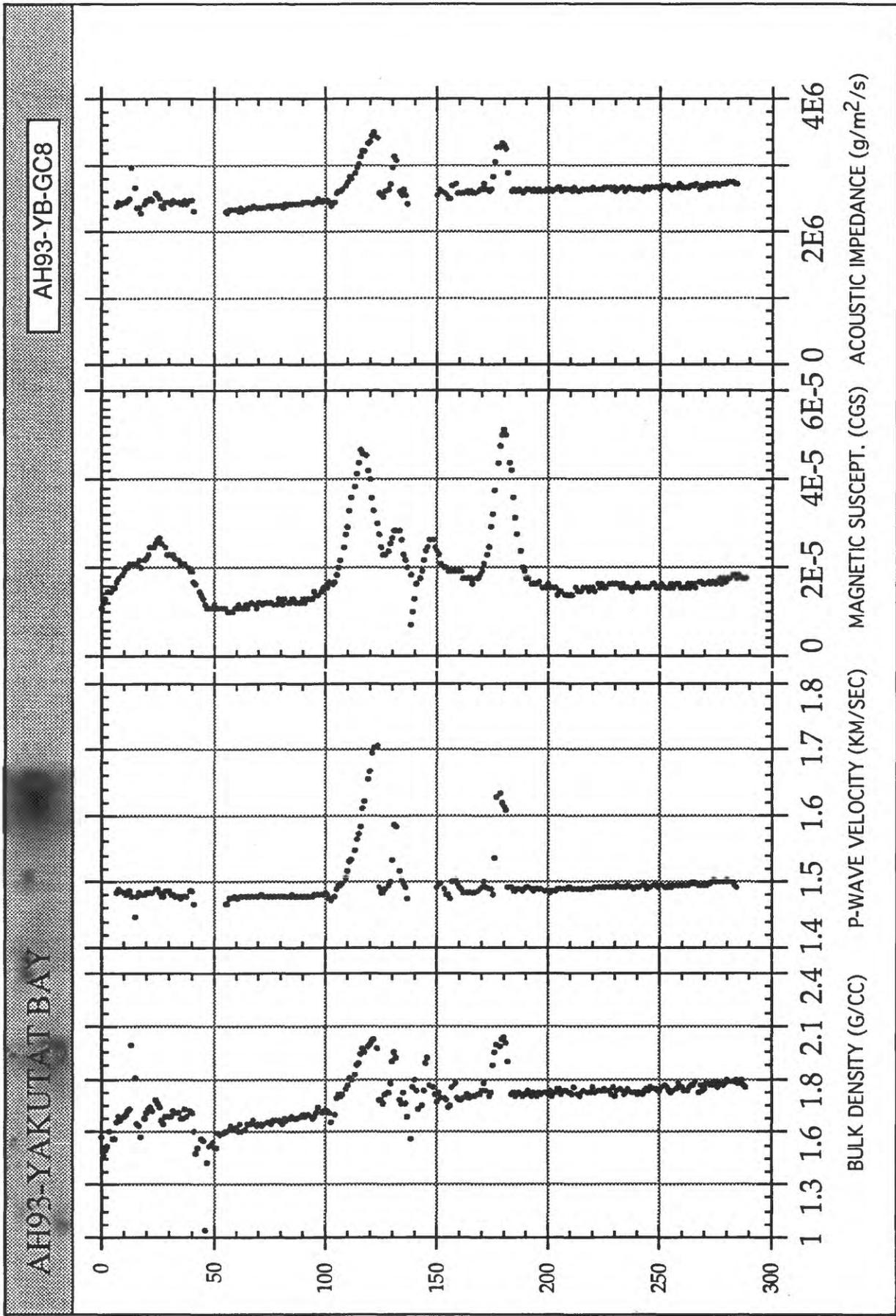




AH93-YAKUTAT BAY

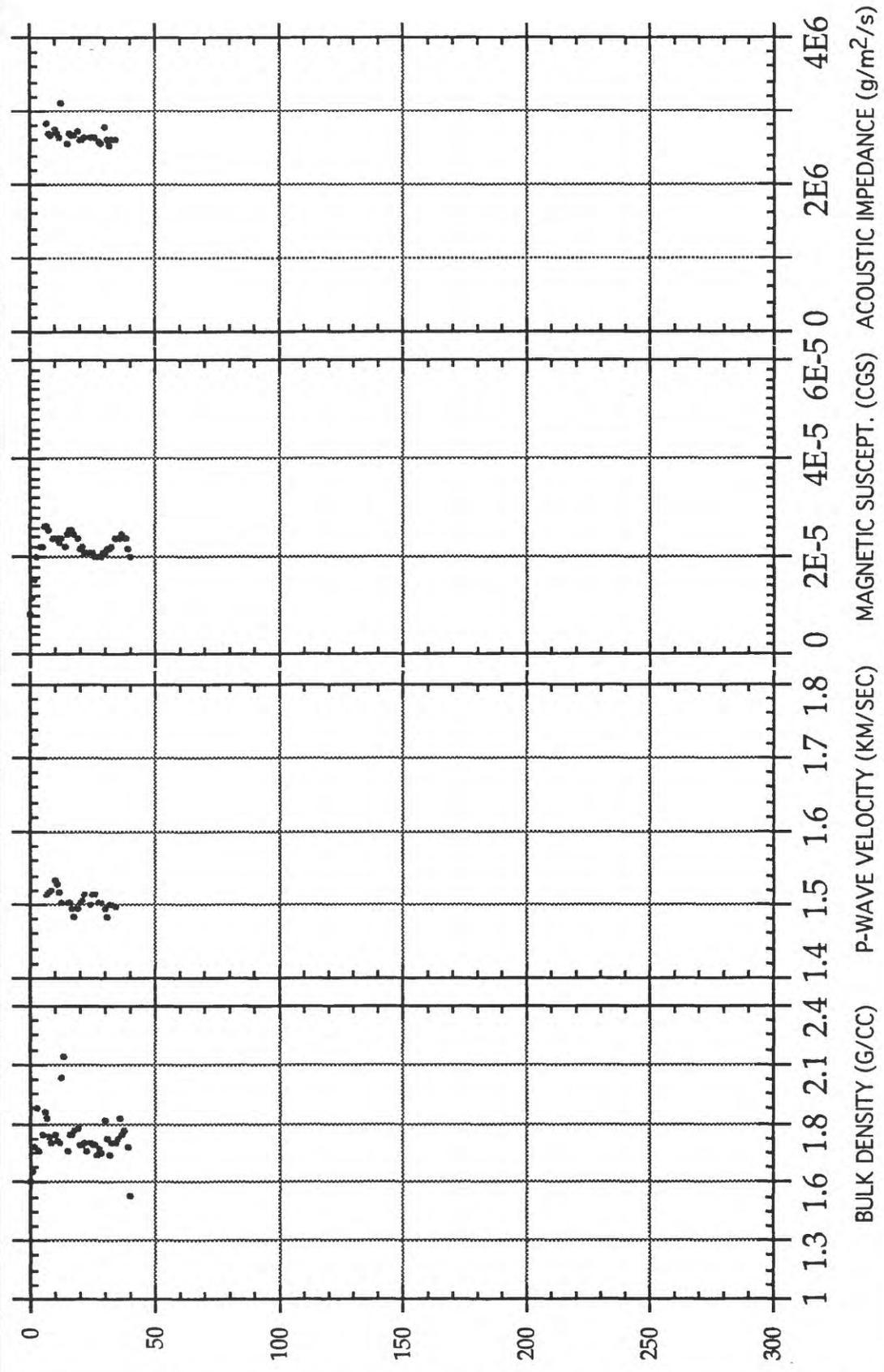
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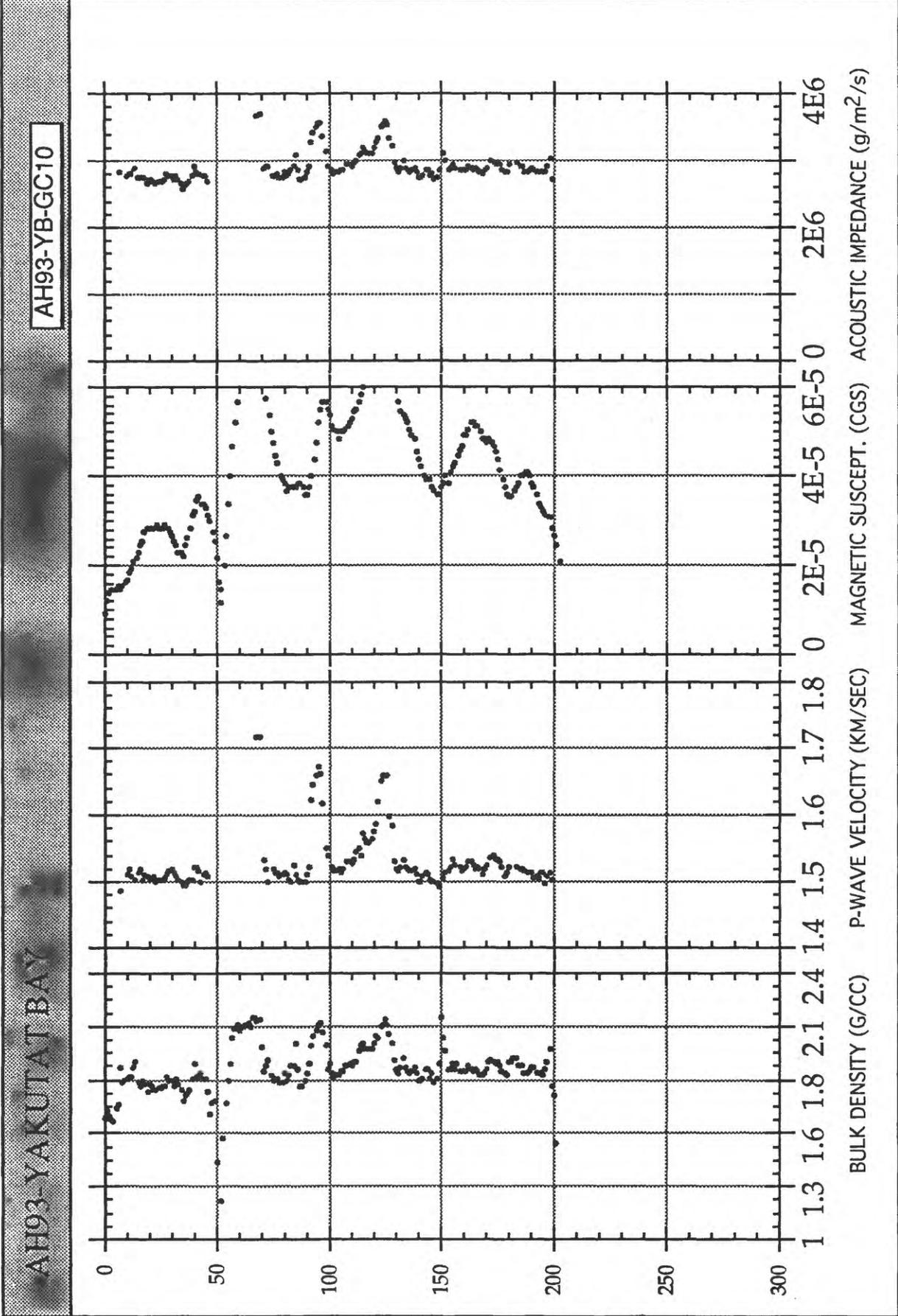


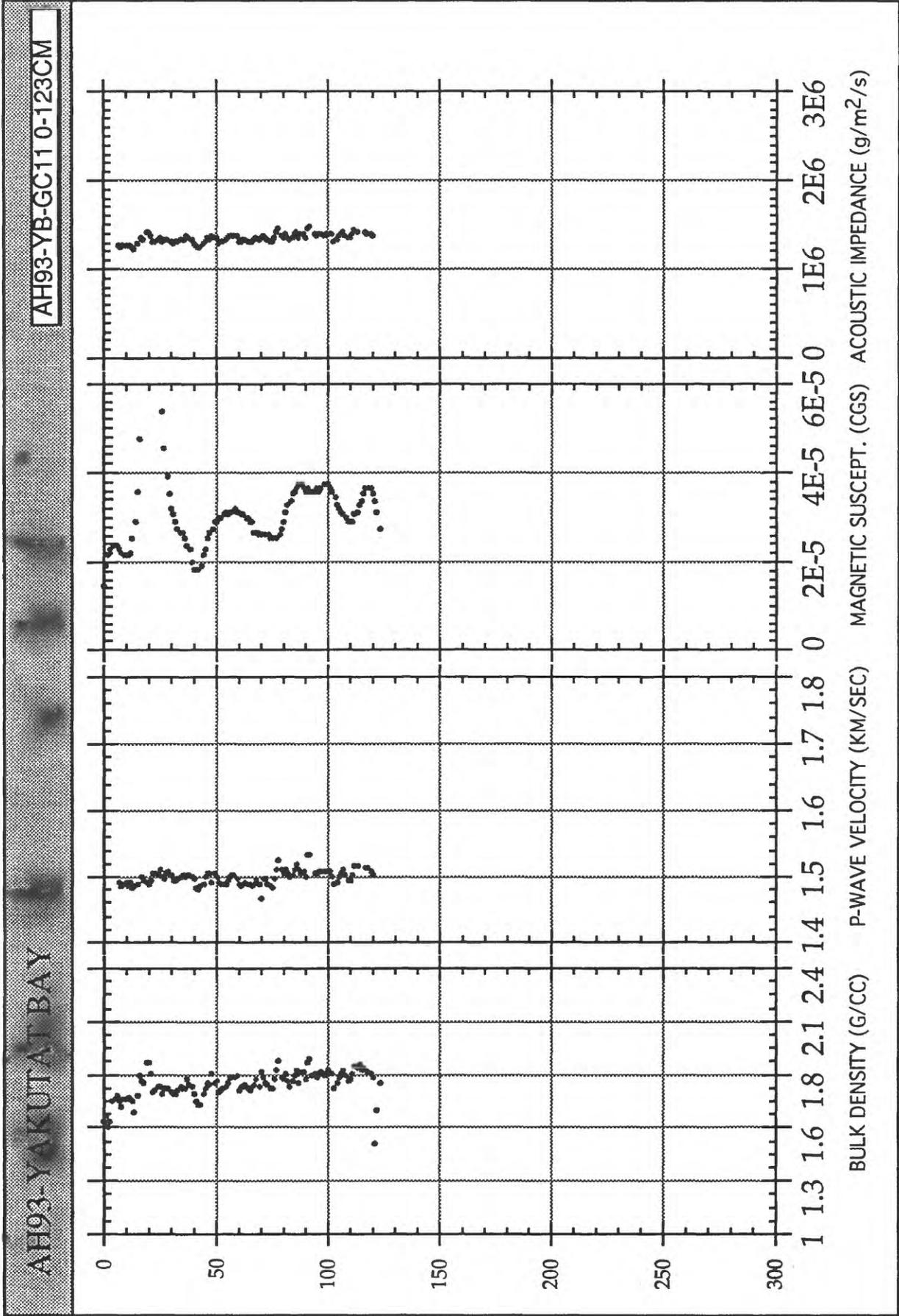


AH93-YAKUTAT BAY

AH93-YB-GC9

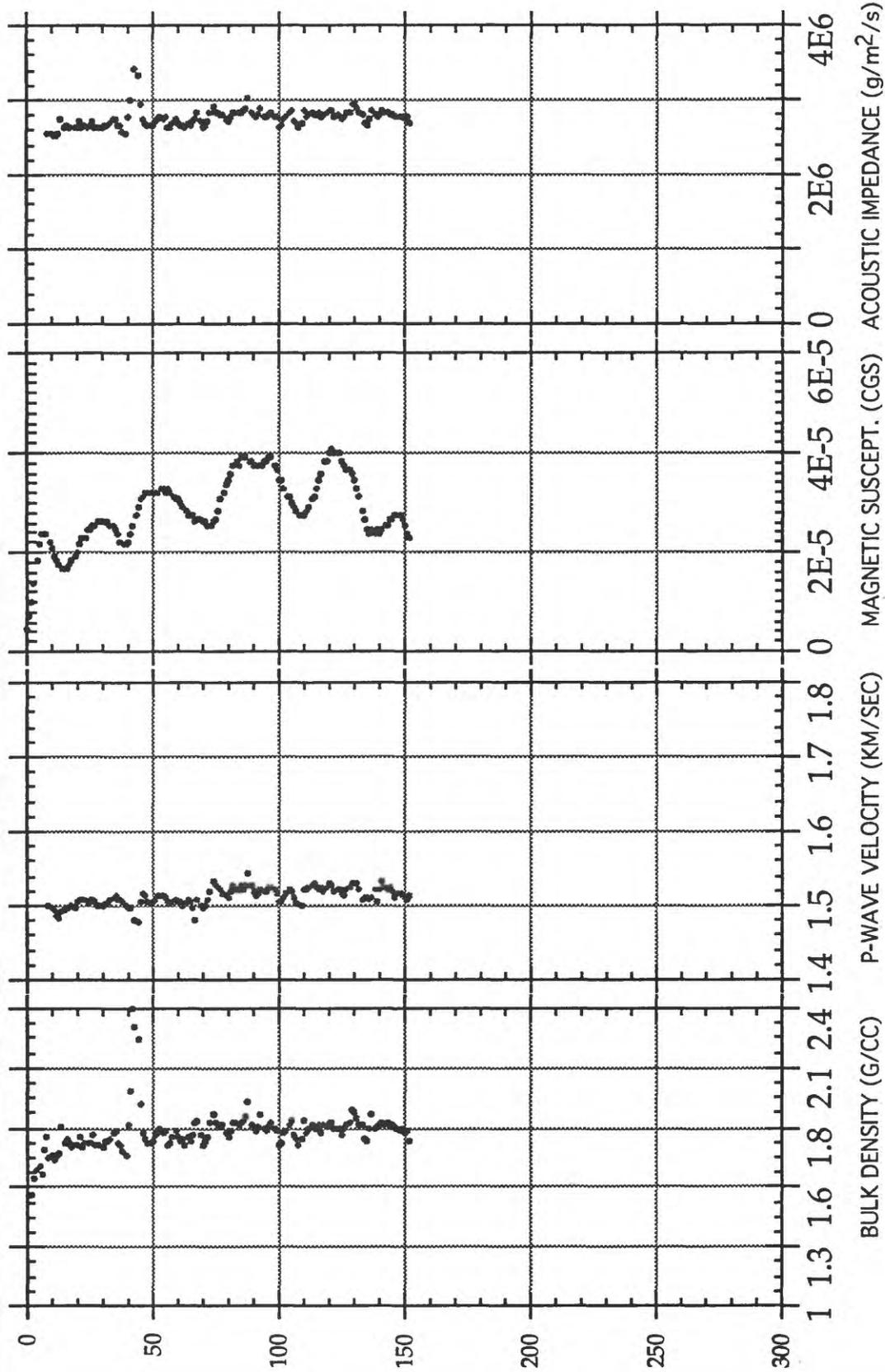


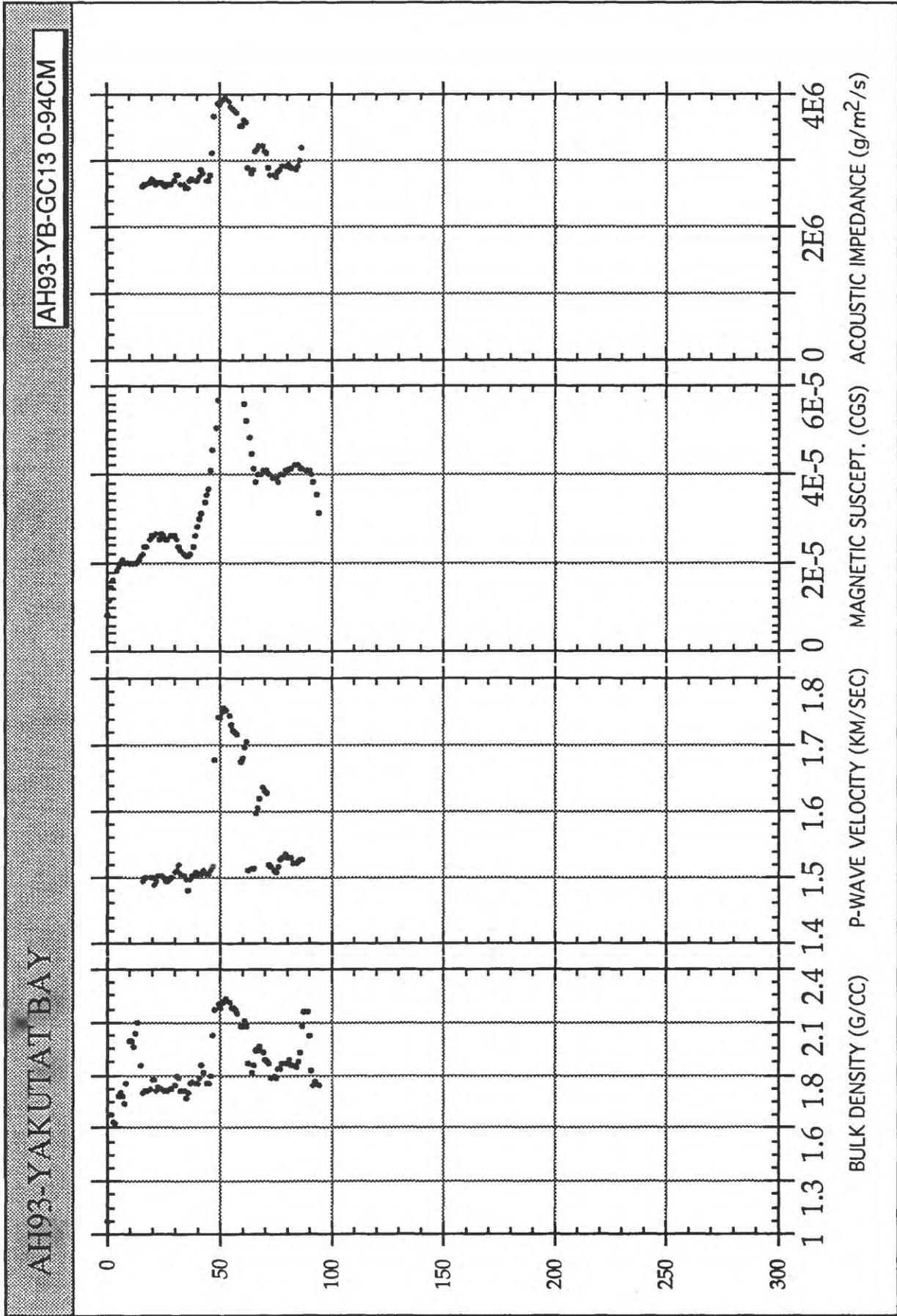


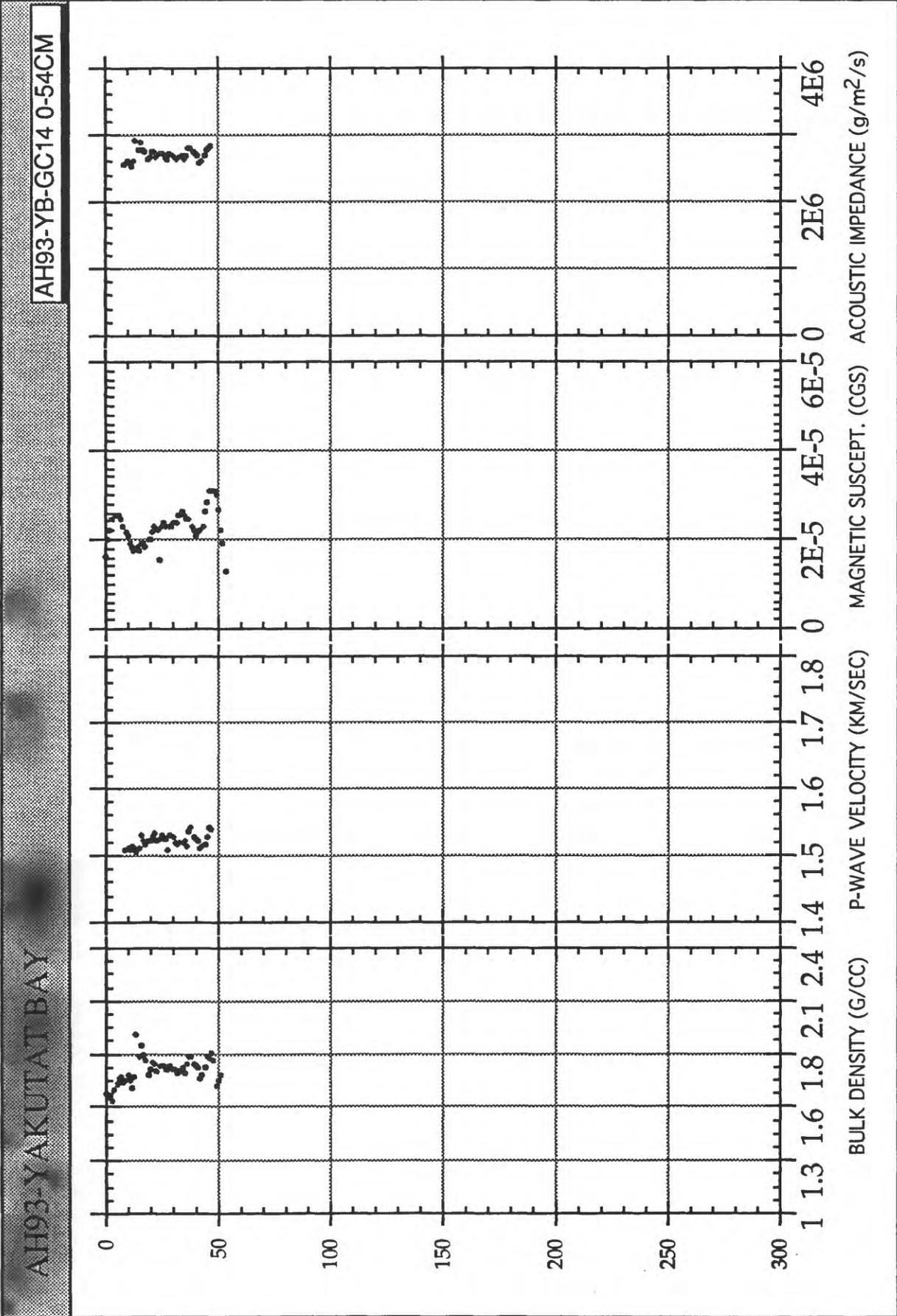


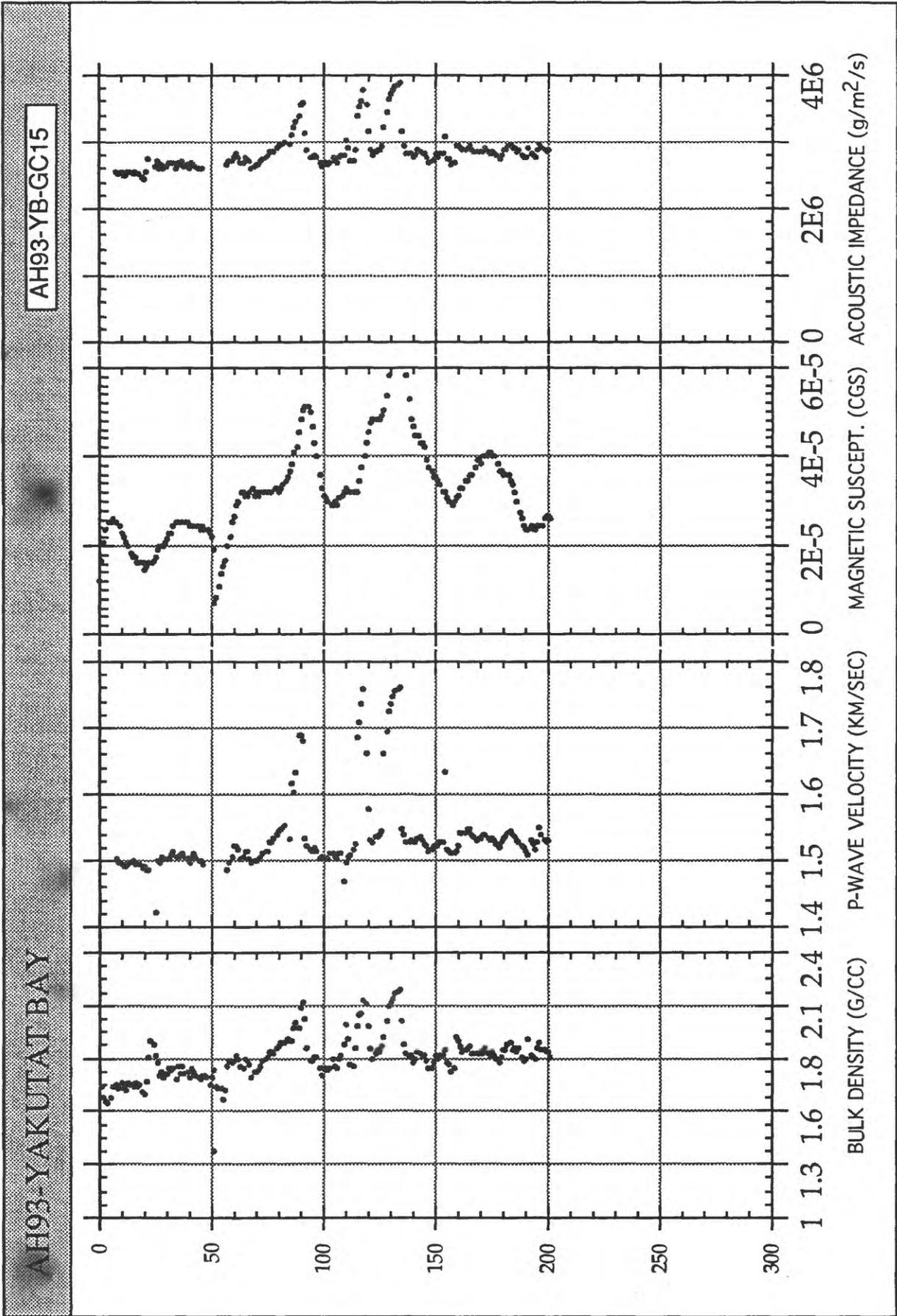
AH93-YAKUTAT BAY

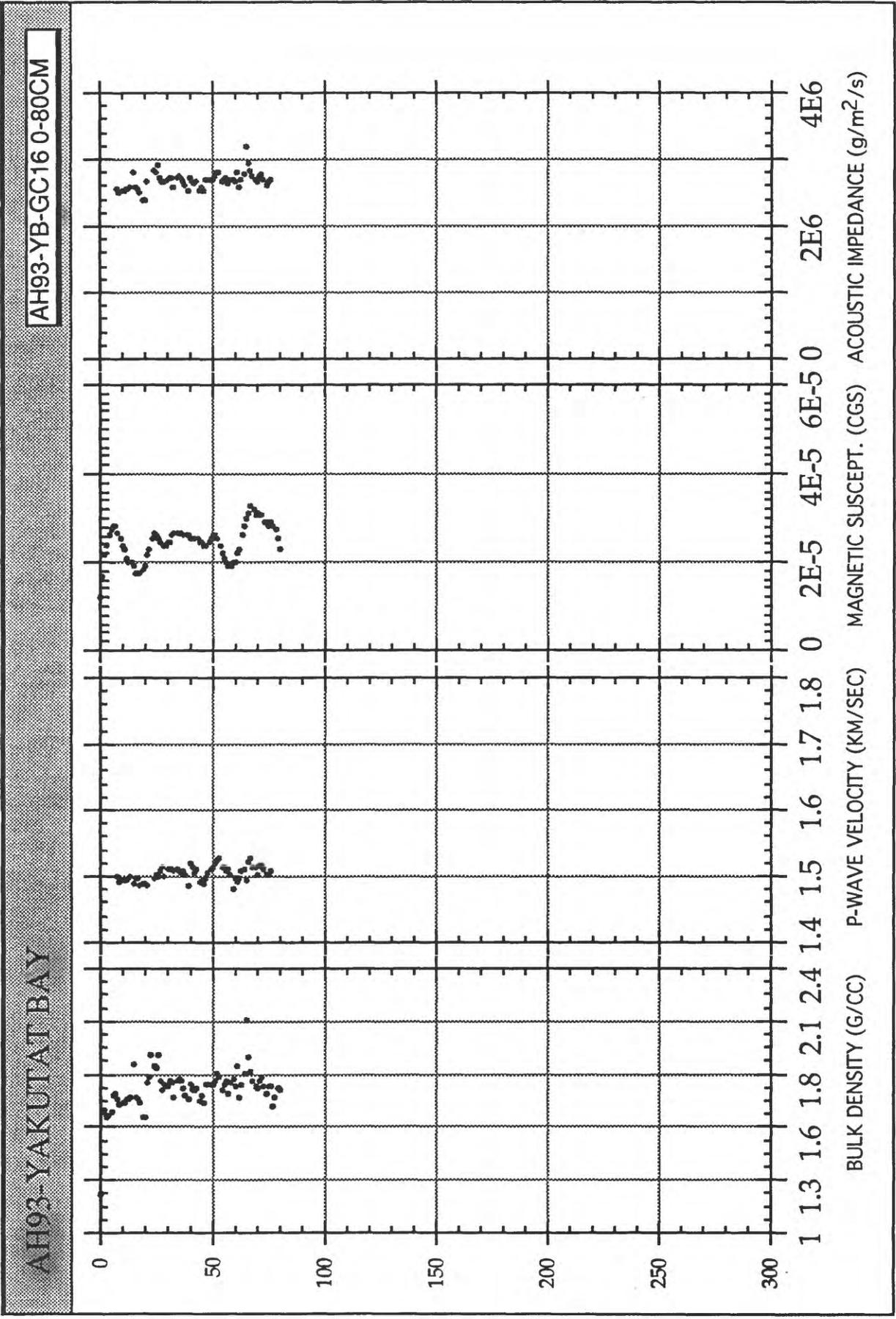
AH93-YB-GC12 0-152

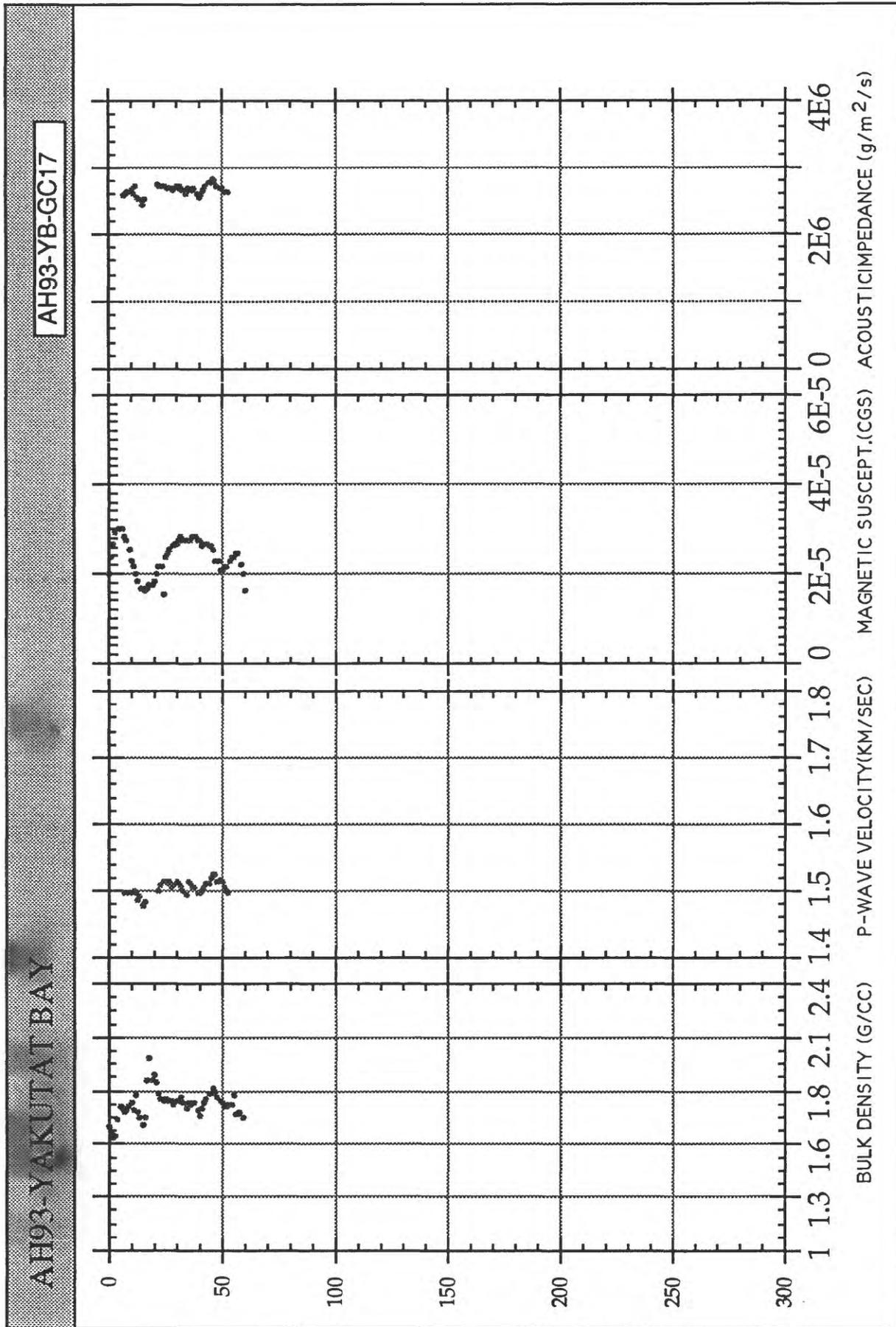






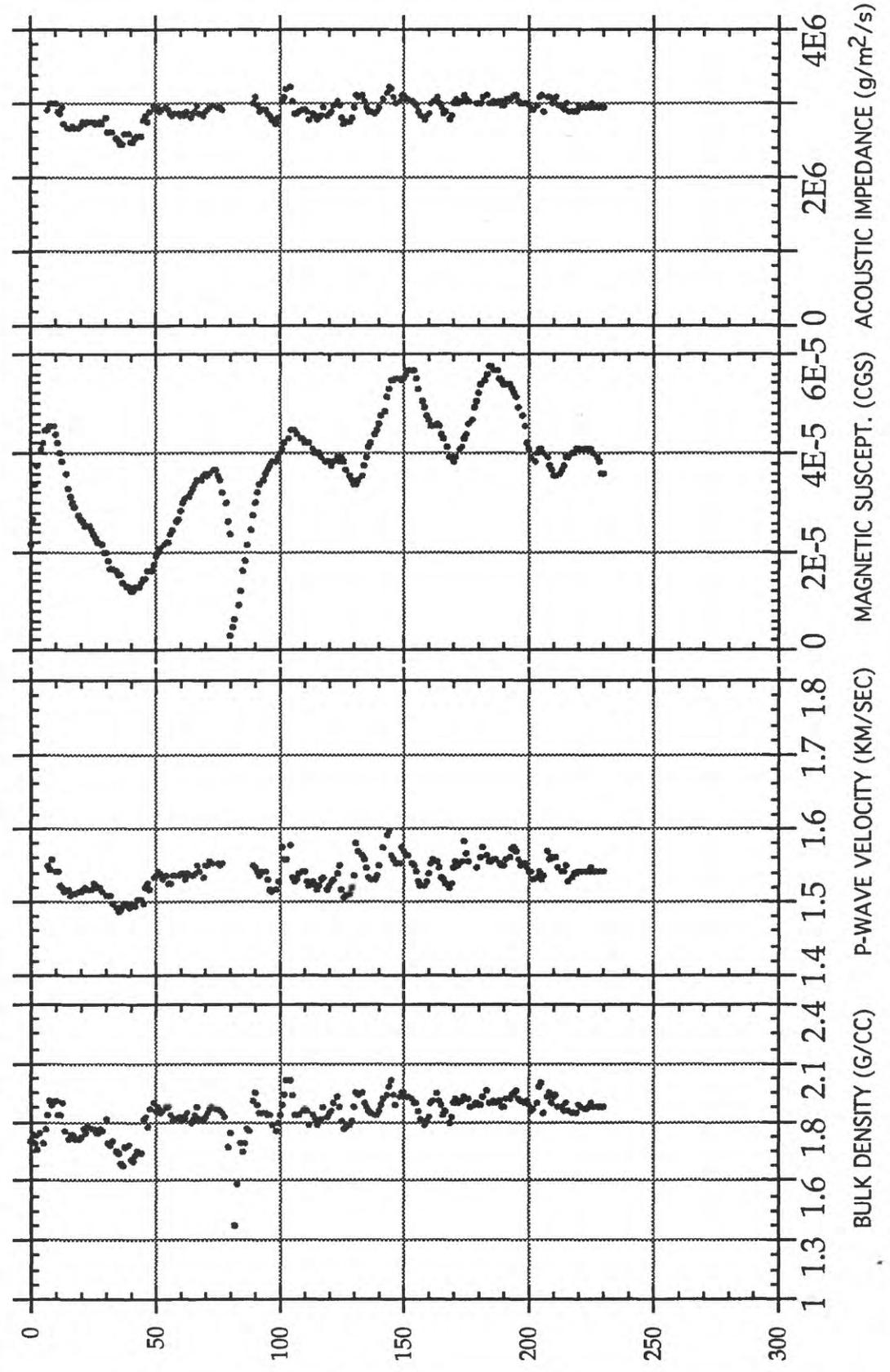


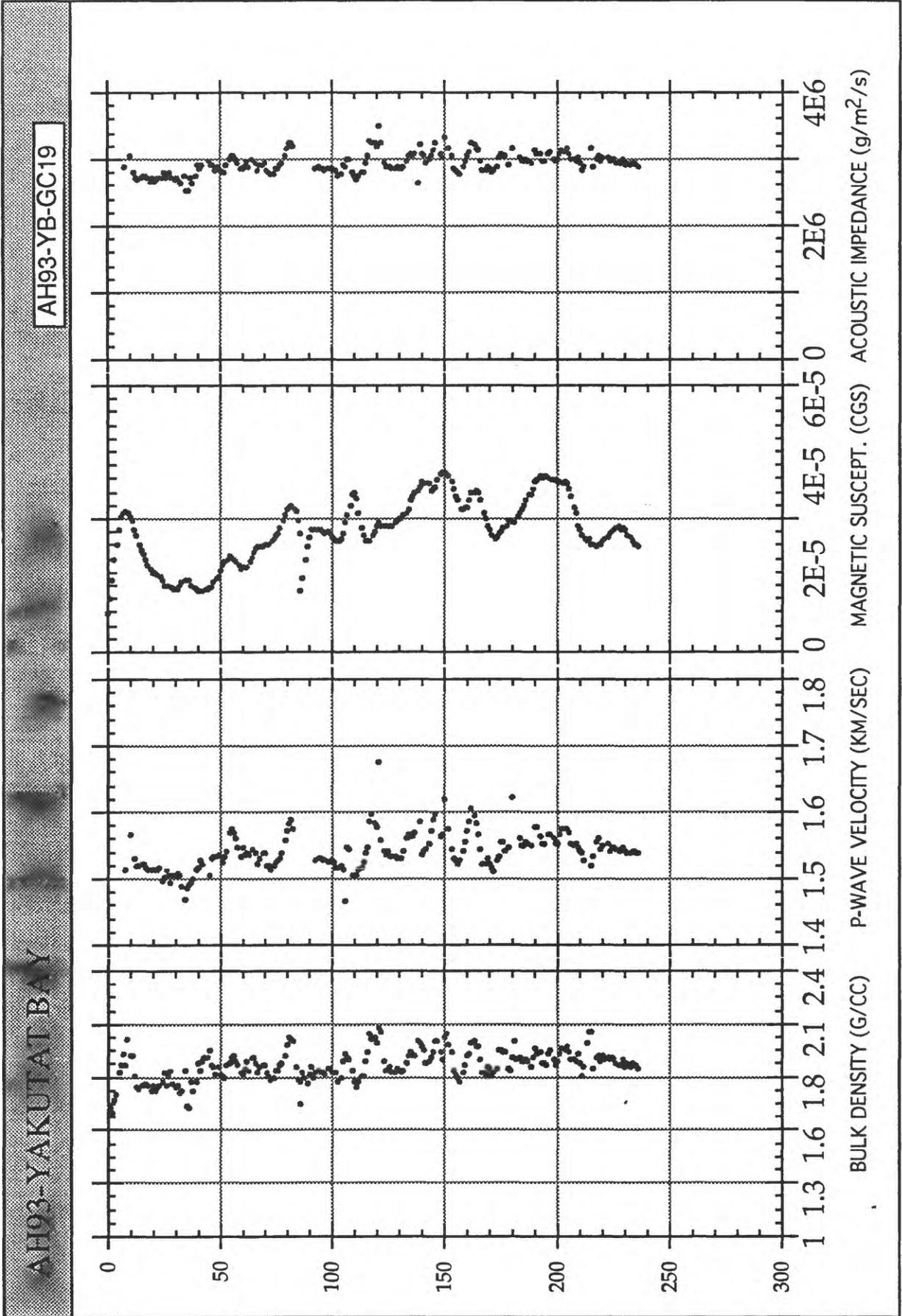


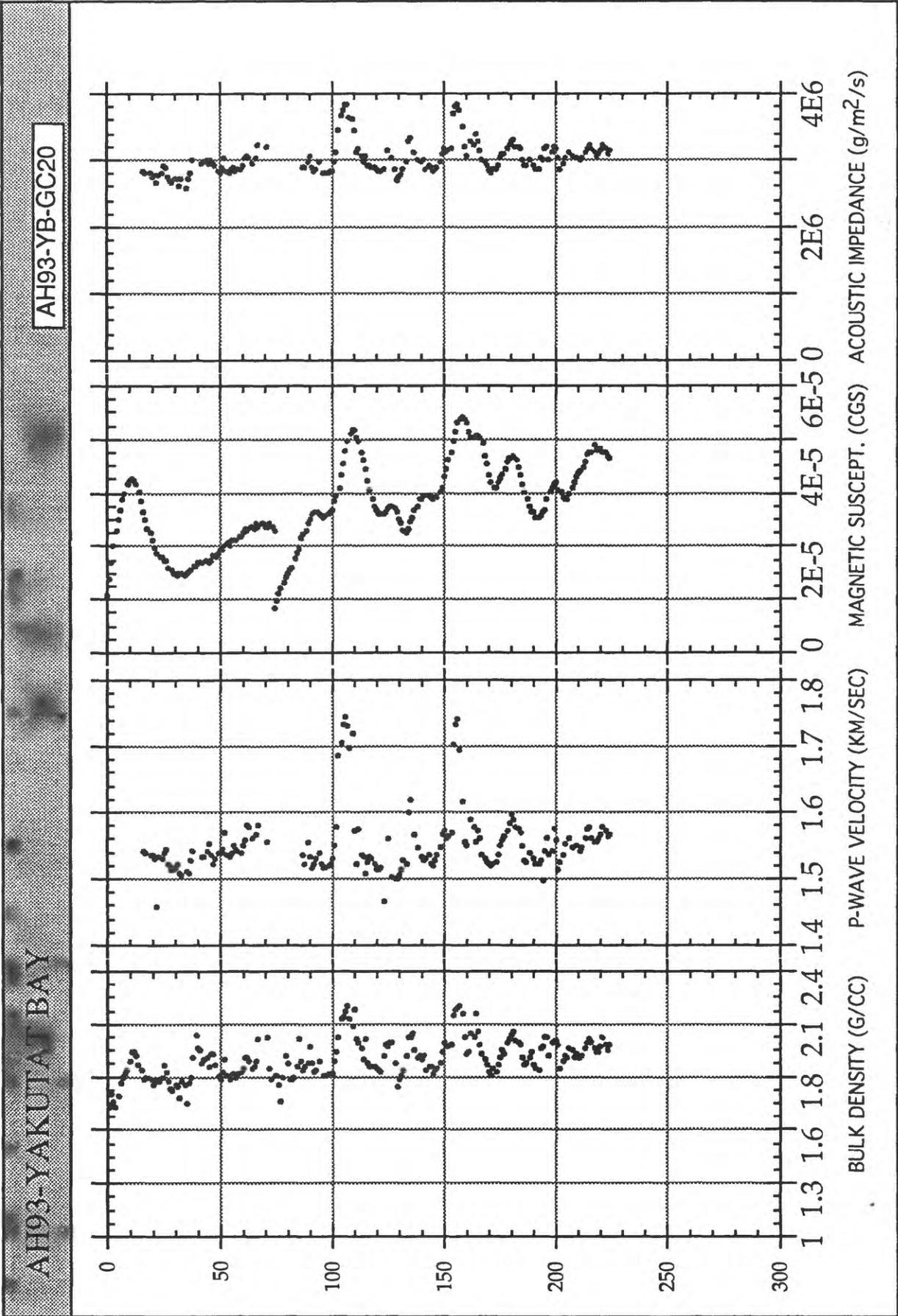


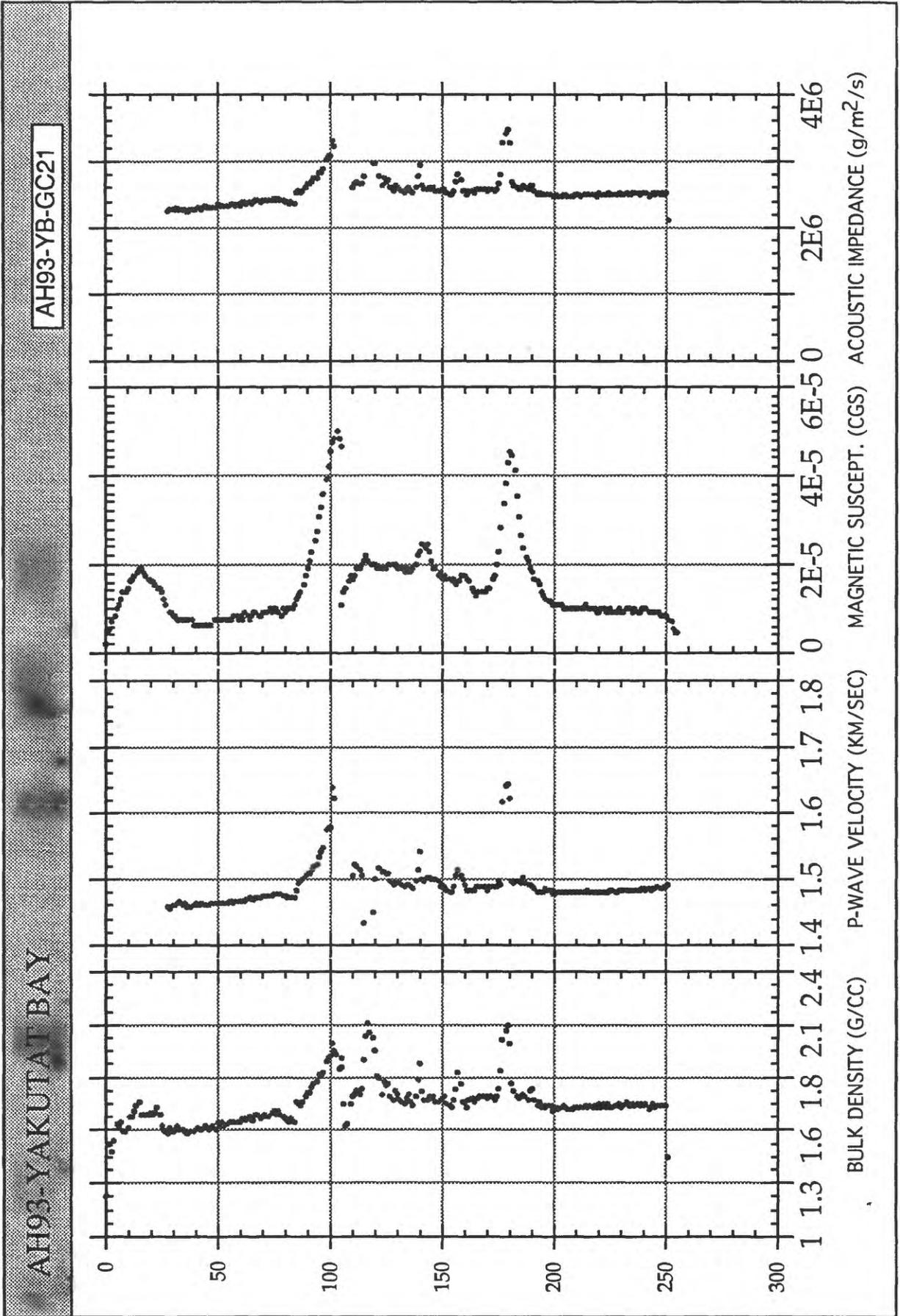
AH93-YAKUTAT BAY

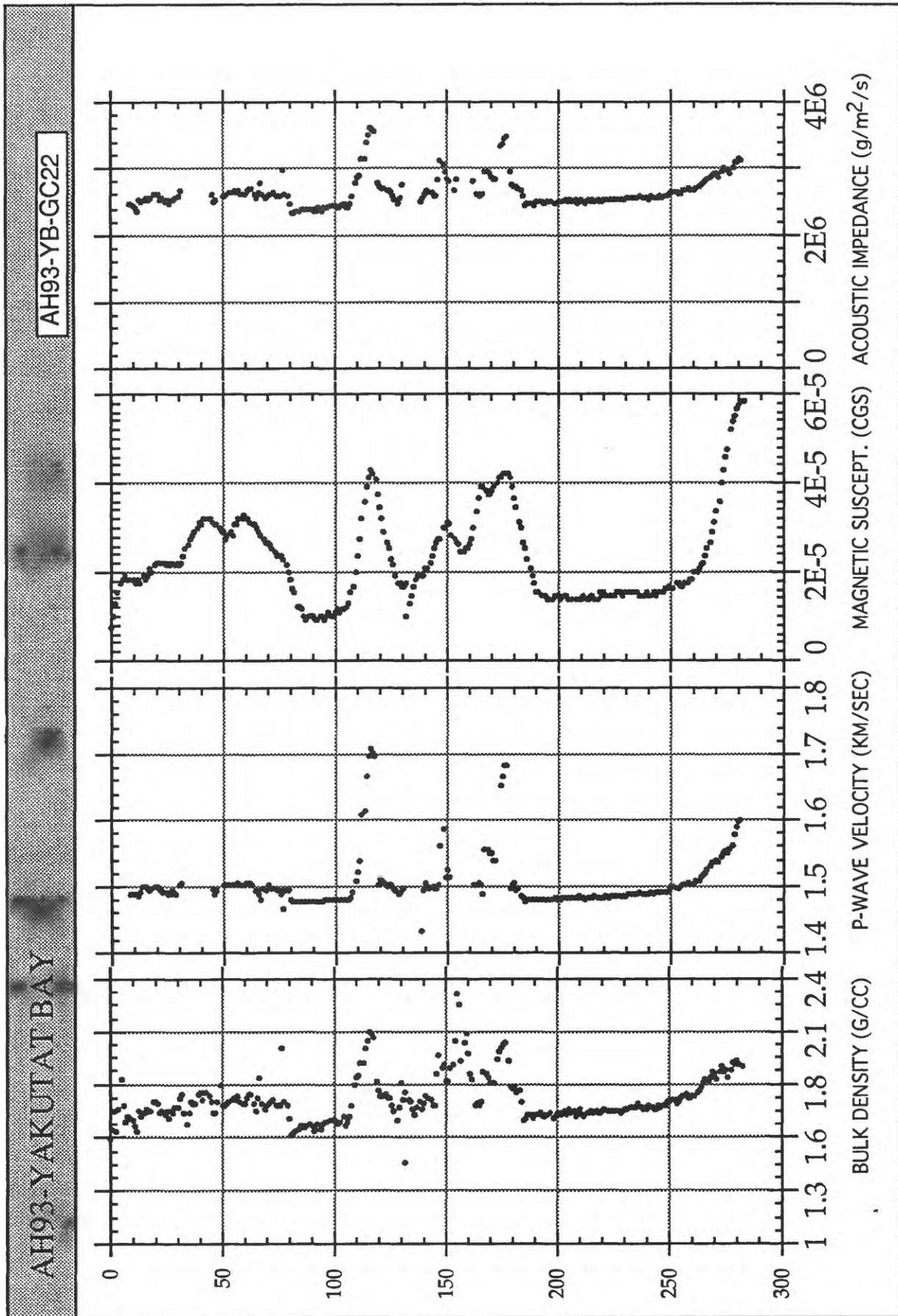
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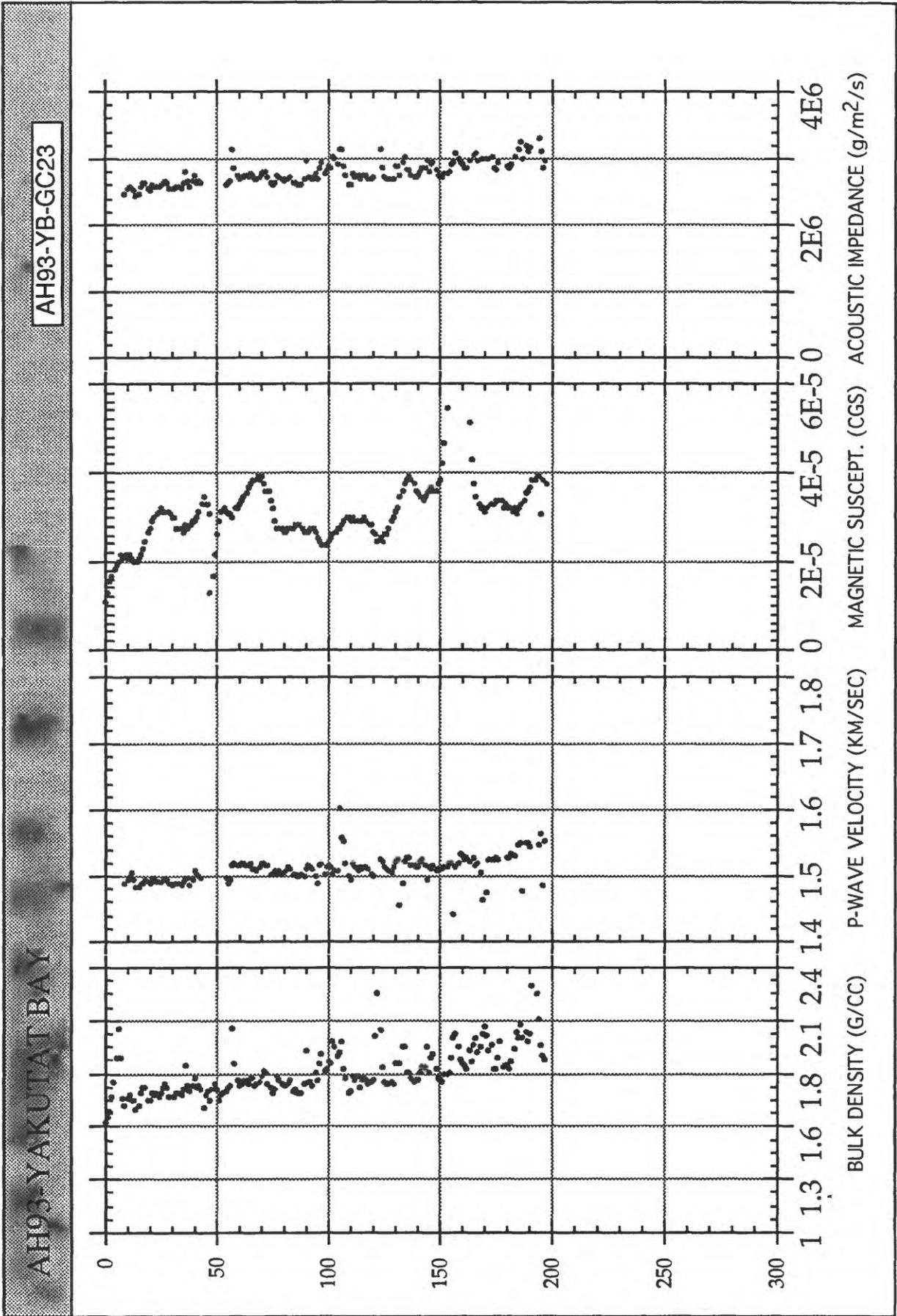


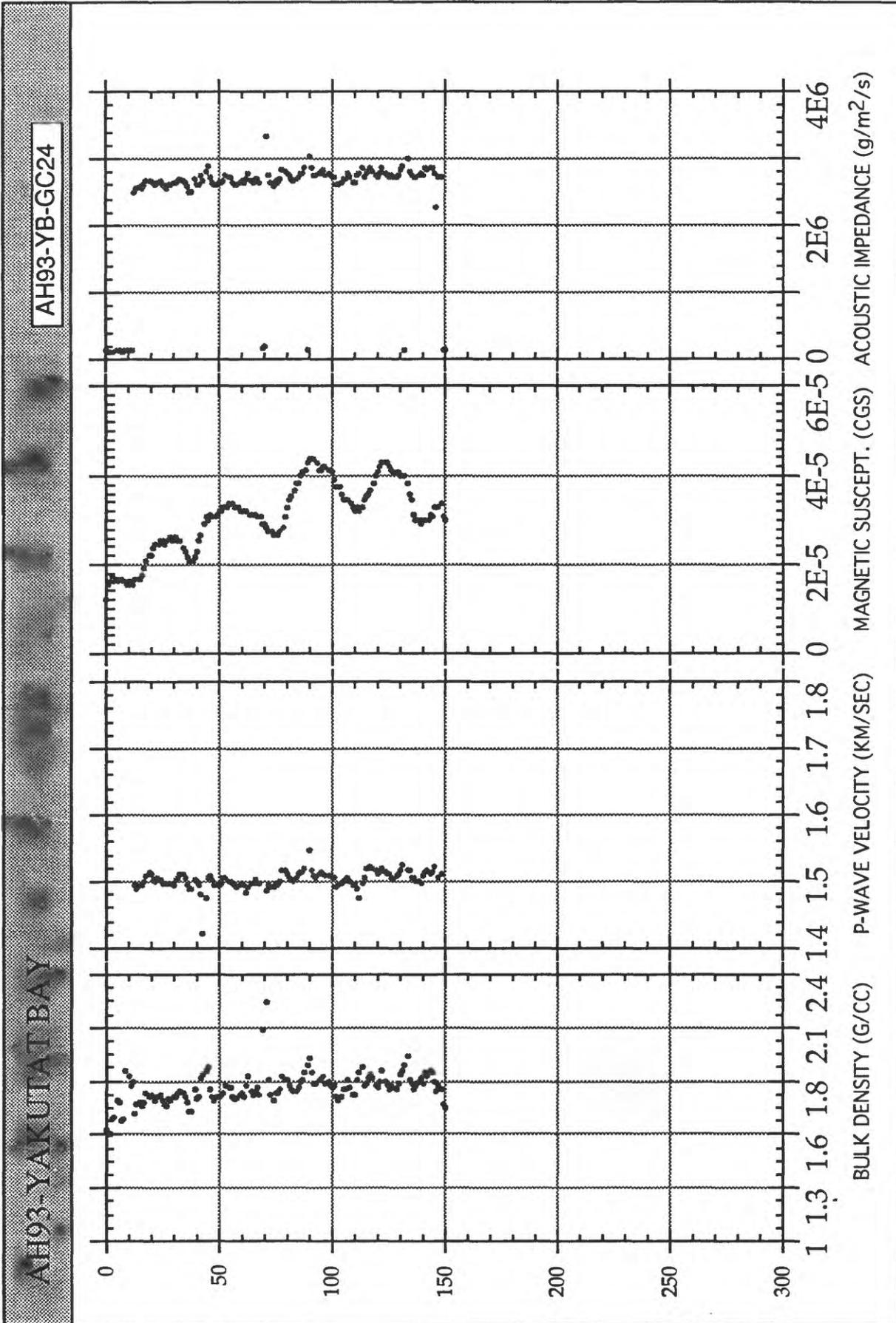






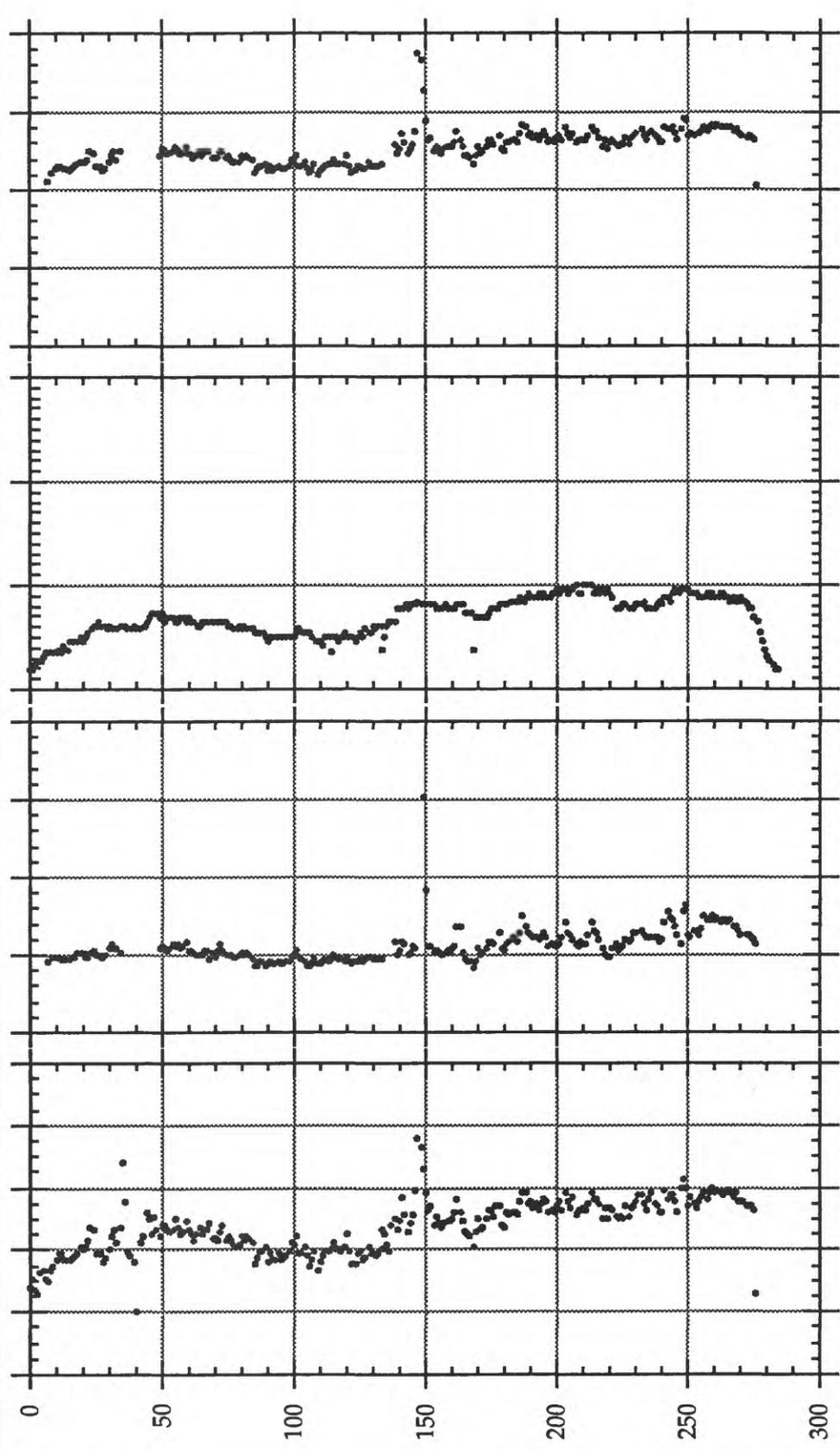






AH93YB-GC25

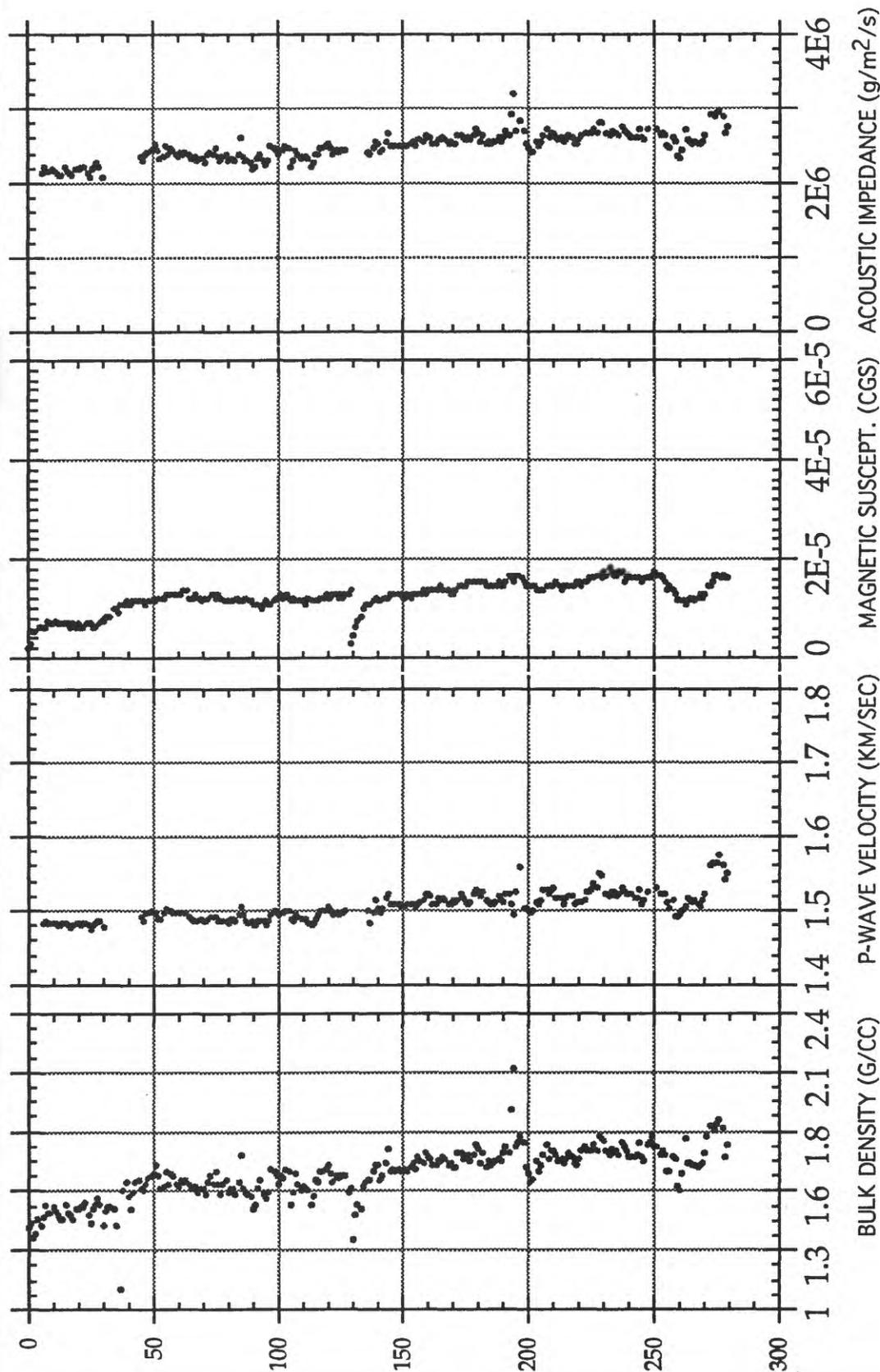
AH93-YAKUTAT BAY

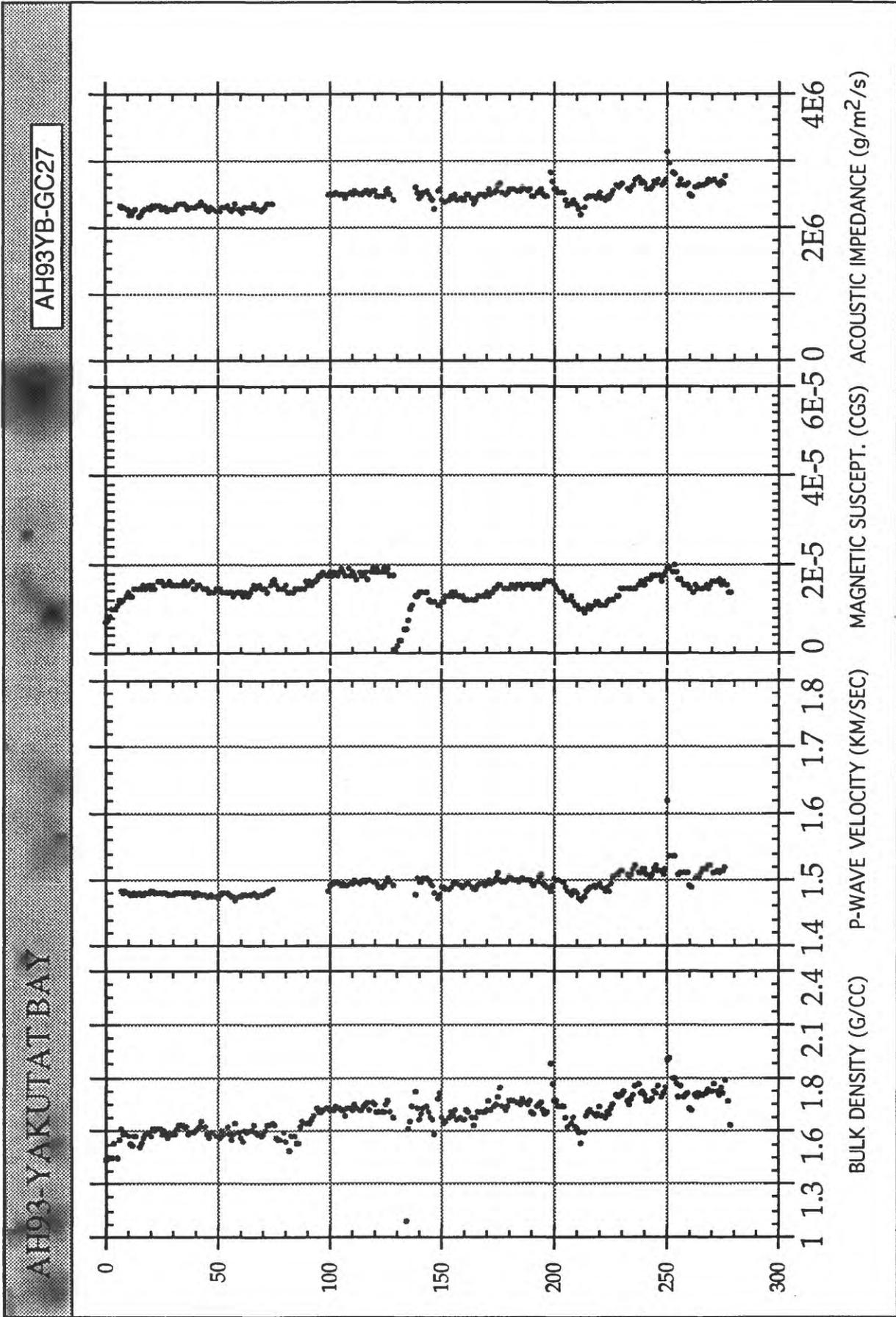


1 1.3 1.6 1.8 2.1 2.4 1.4 1.5 1.6 1.7 1.8 0 2E-5 4E-5 6E-5 0 2E6 4E6
BULK DENSITY (G/CC) P-WAVE VELOCITY (KM/SEC) MAGNETIC SUSCEPT. (CGS) ACOUSTIC IMPEDANCE (g/m²/s)

AH93-YAKUTAI BAY

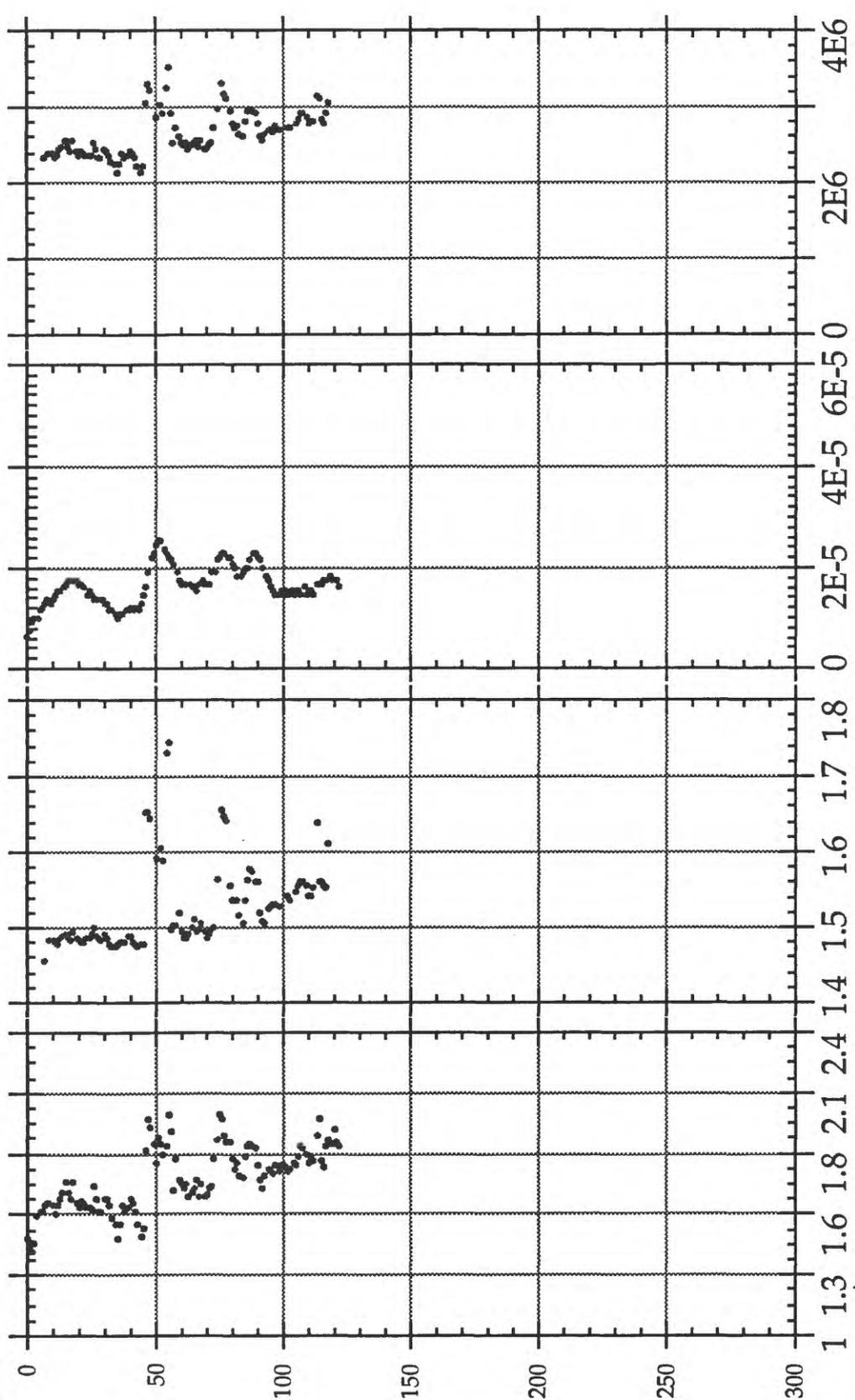
AH93-YB-GC26





AH93-YAKUTAT BAY

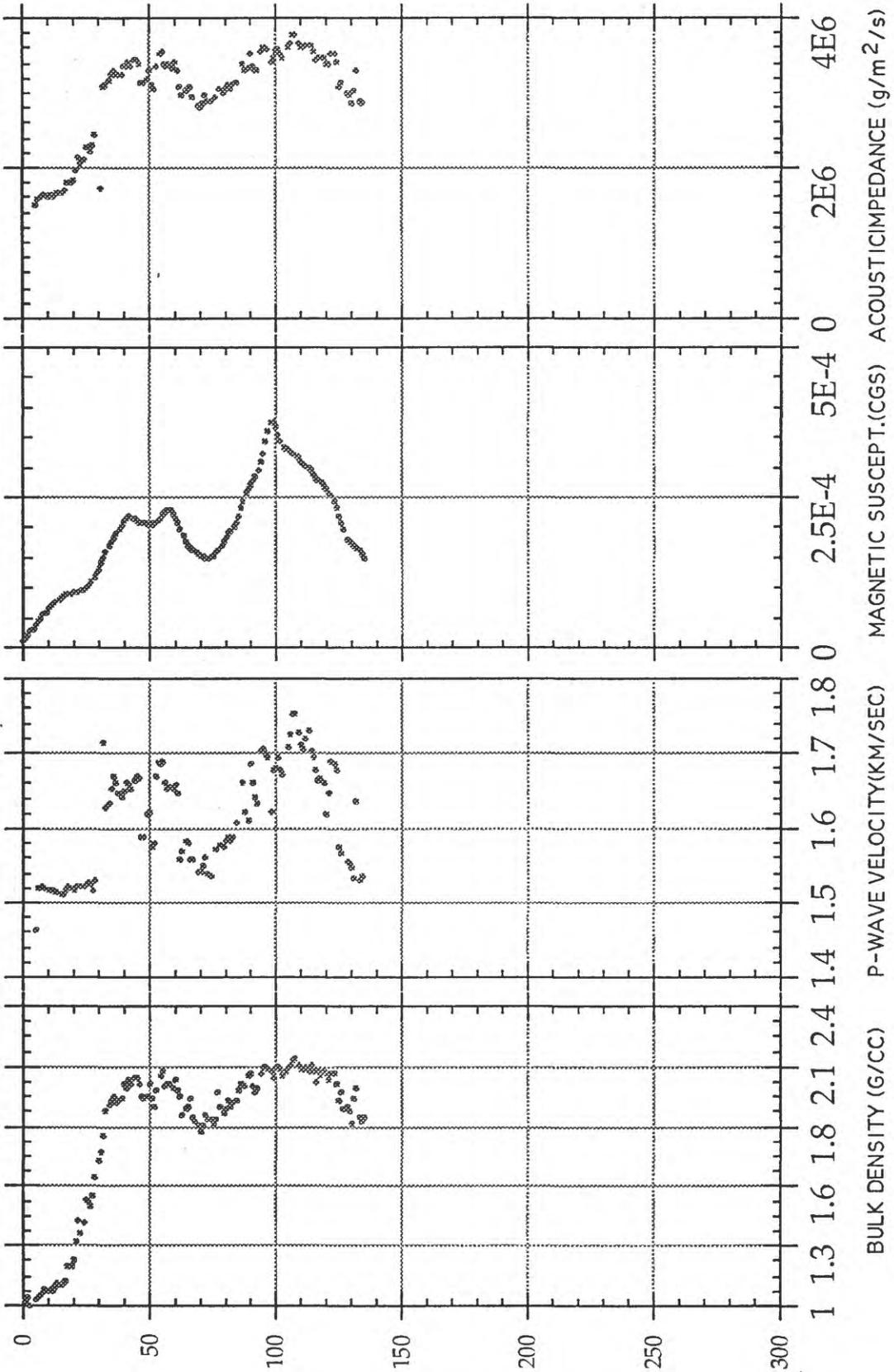
AH93-YB-GC28



BULK DENSITY (G/CC) P-WAVE VELOCITY (KM/SEC) MAGNETIC SUSCEPT. (CGS) ACOUSTIC IMPEDANCE (G/M²/S)

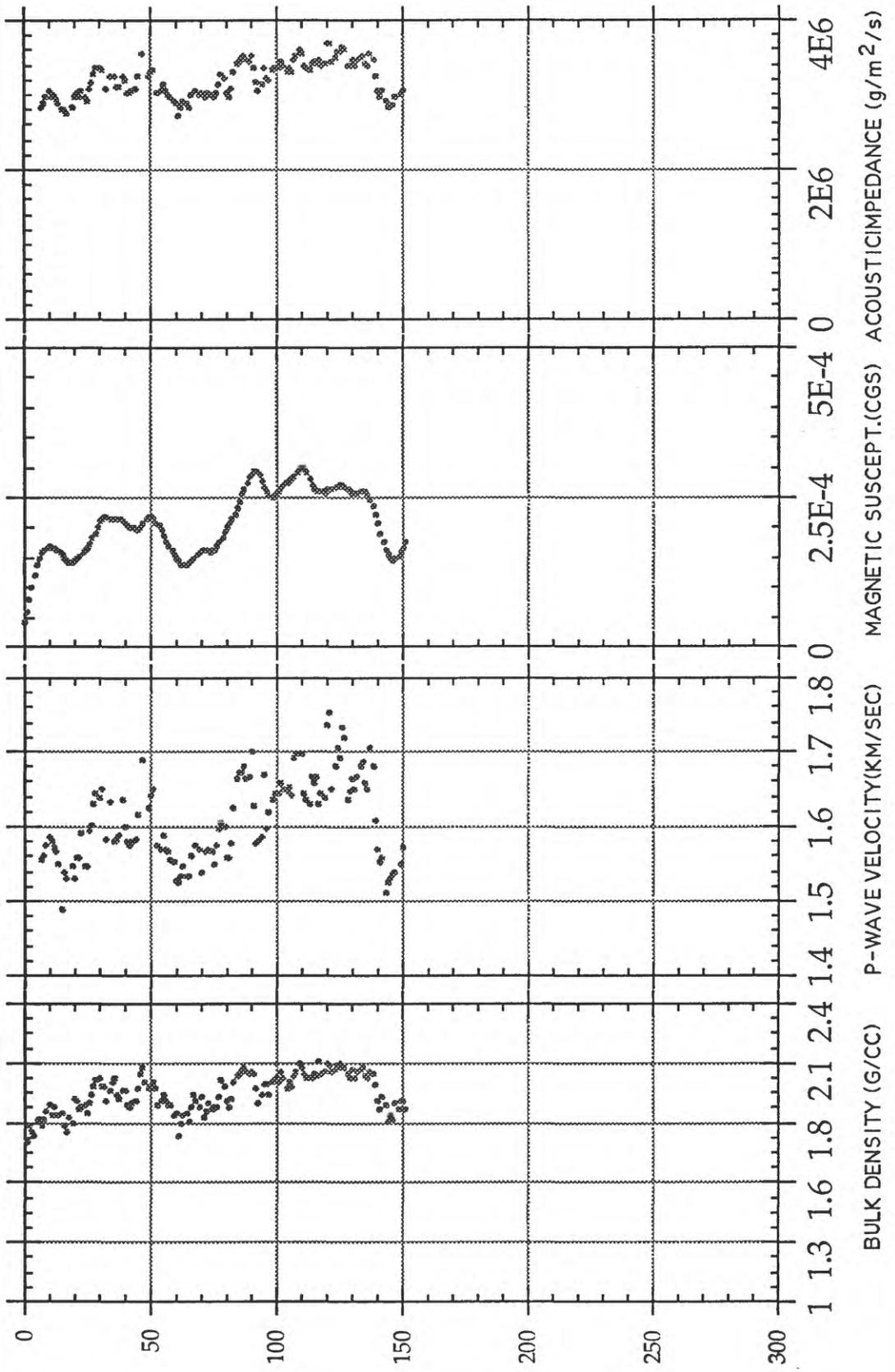
AH93-GLACIER BAY

AH93-GB-GC2 0-135



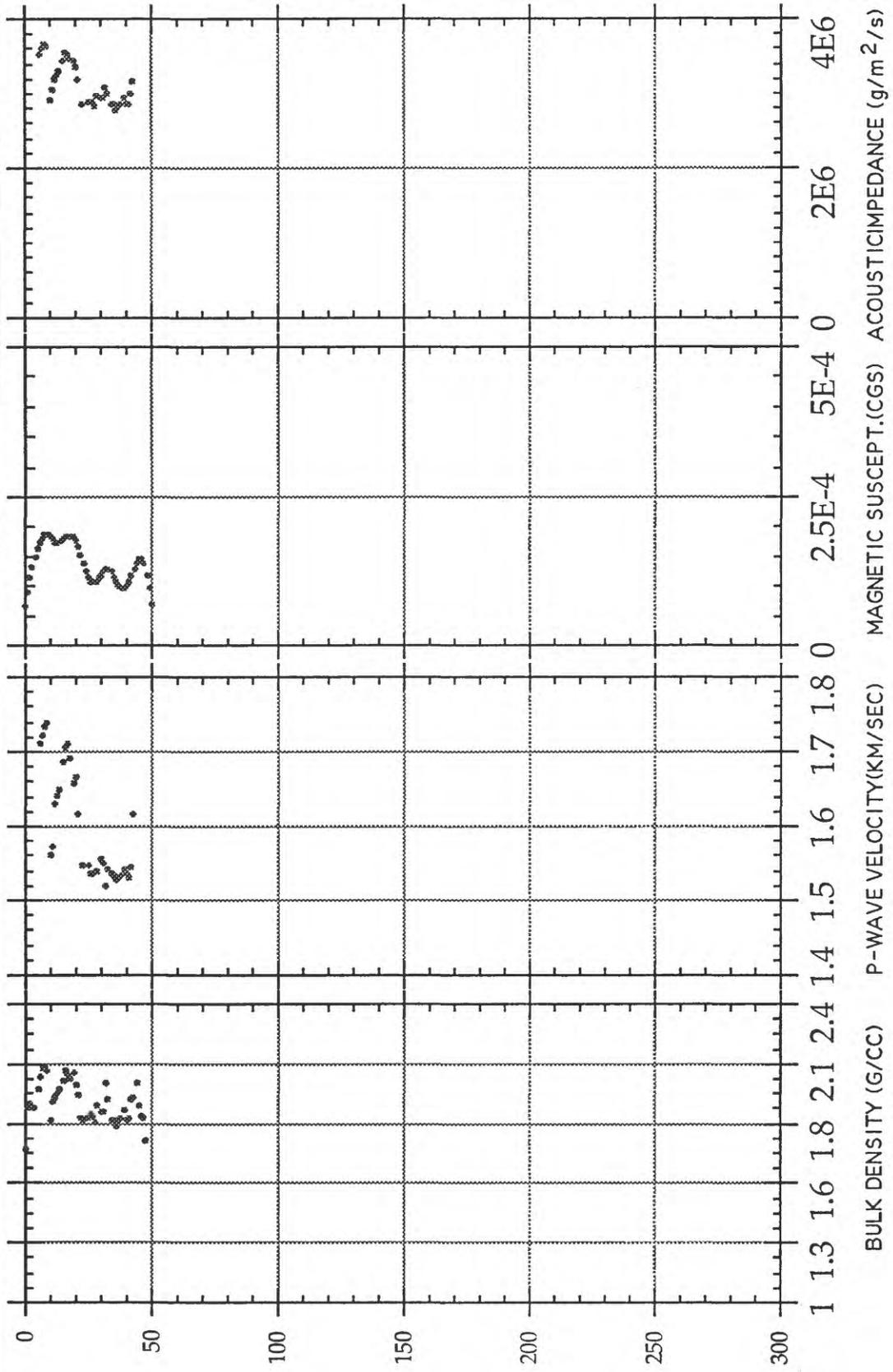
AH93-GLACIER BAY

AH93-GB-GC3 0-151



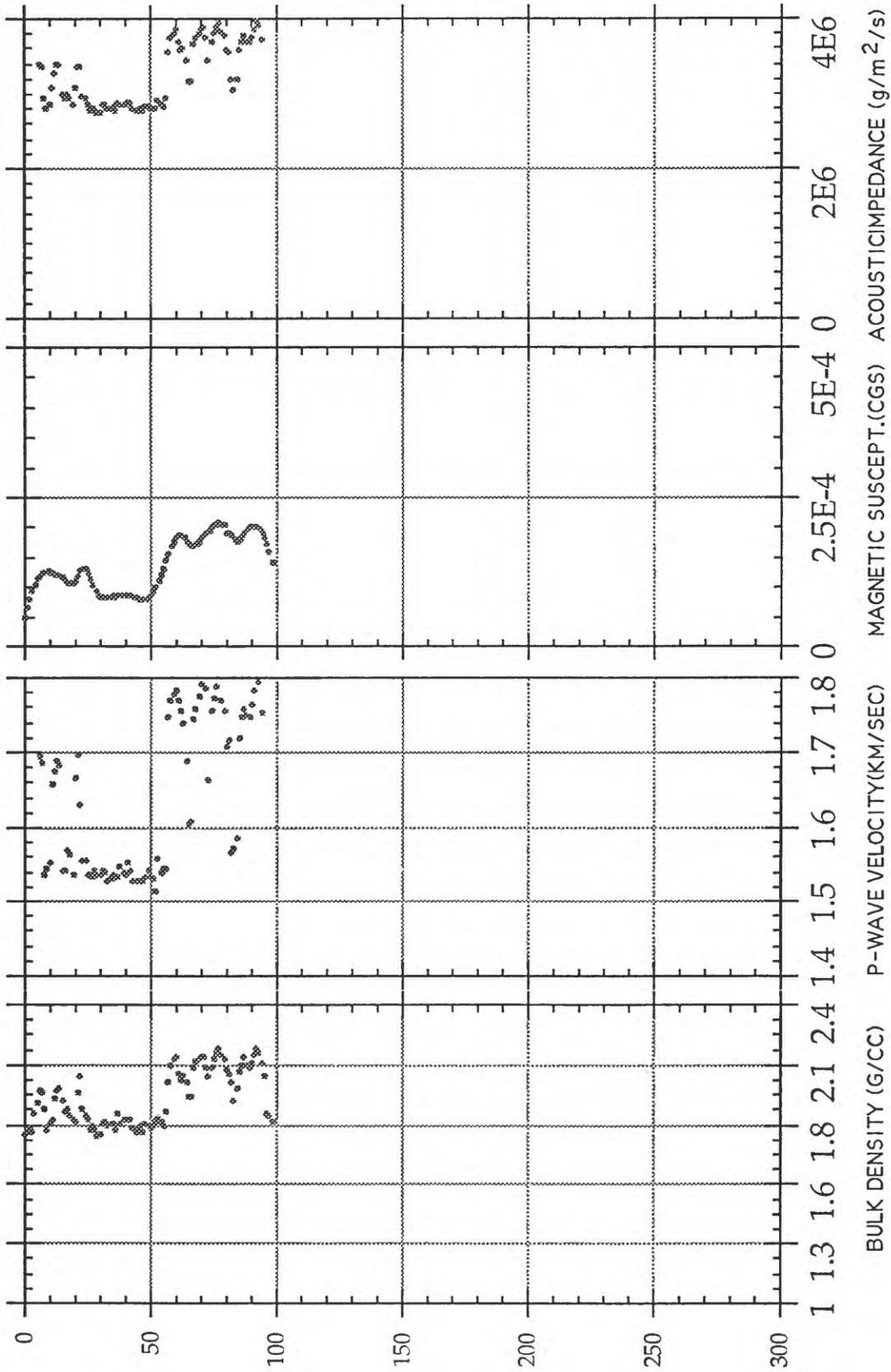
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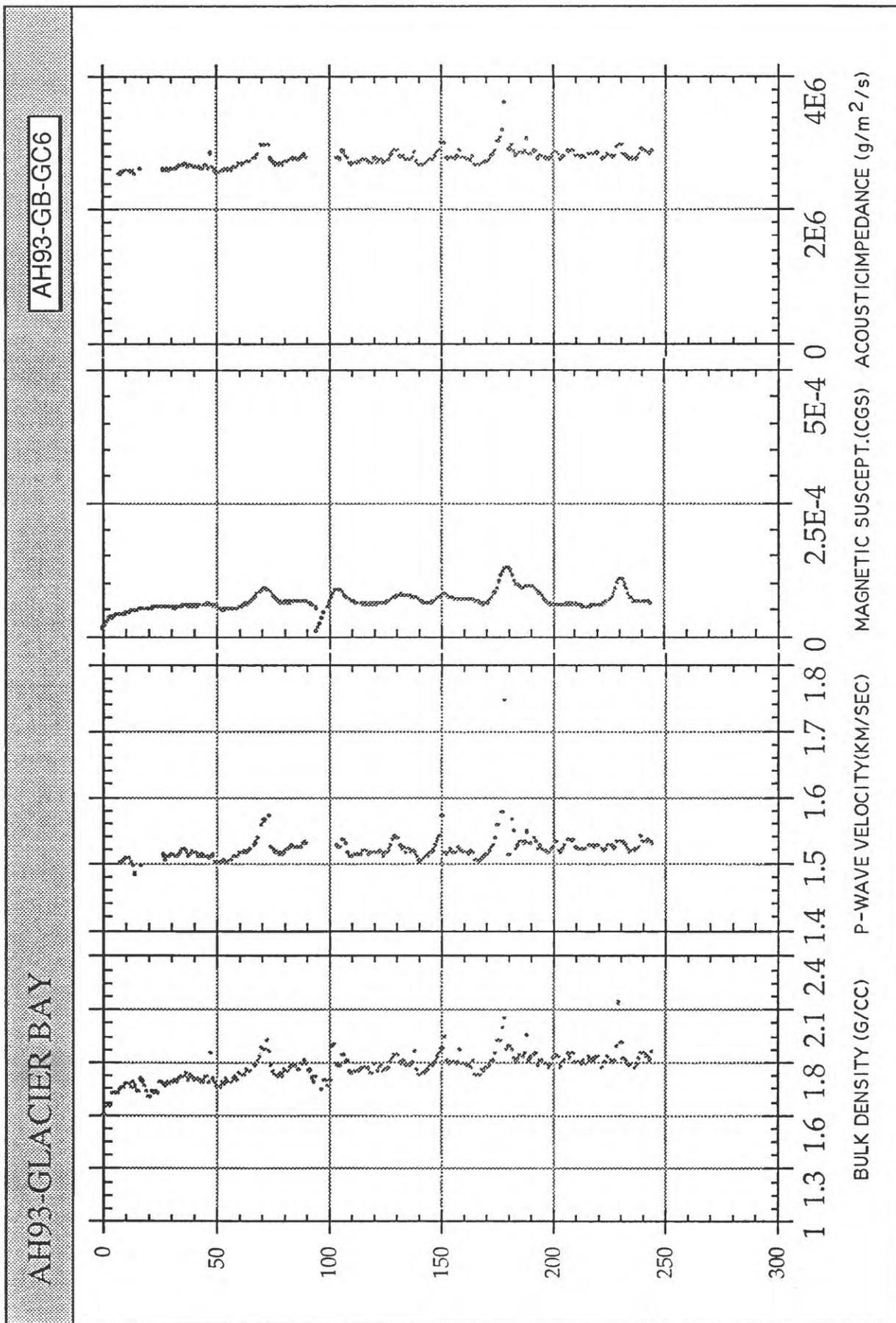
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AH93-GLACIER BAY

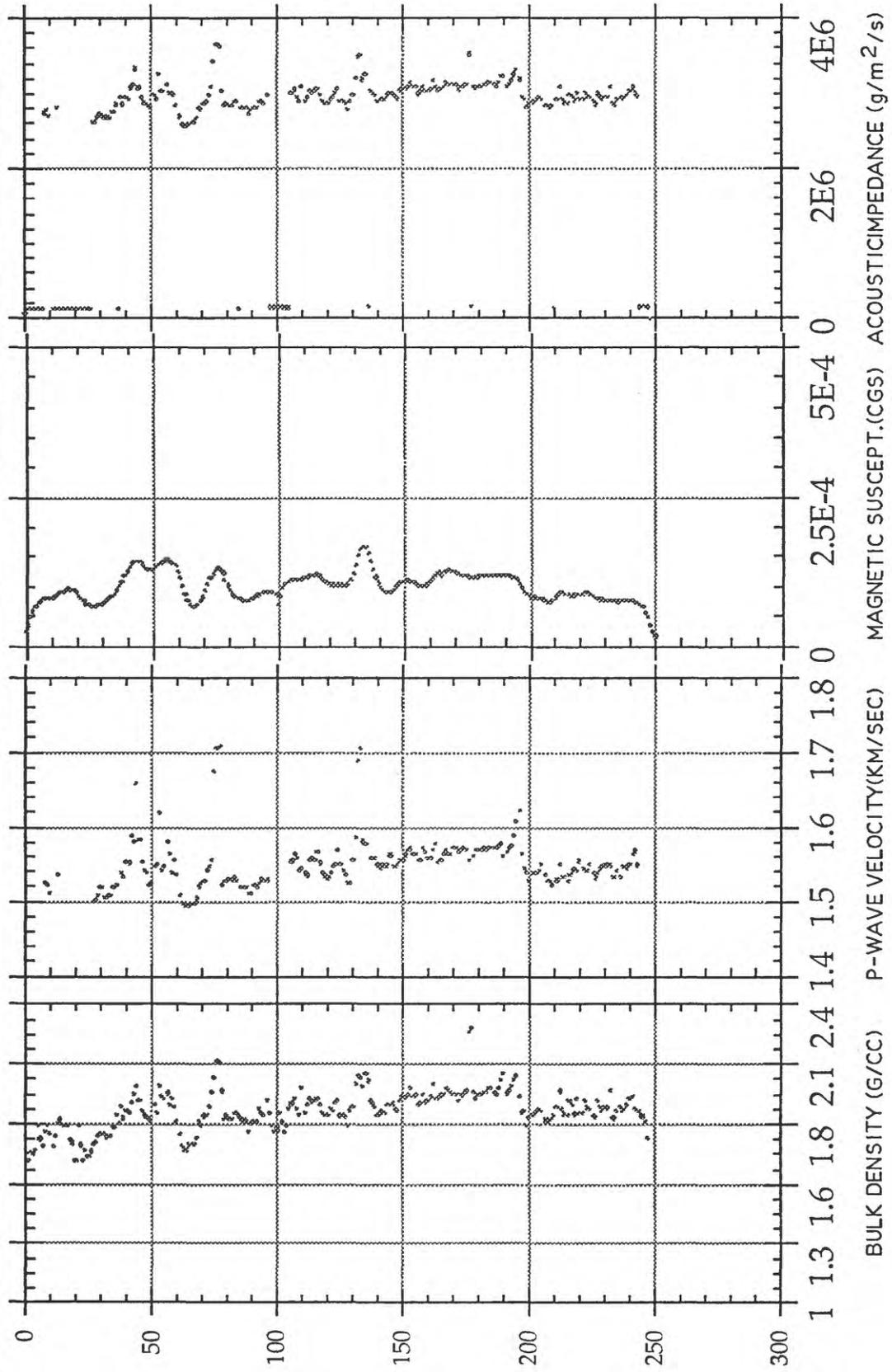
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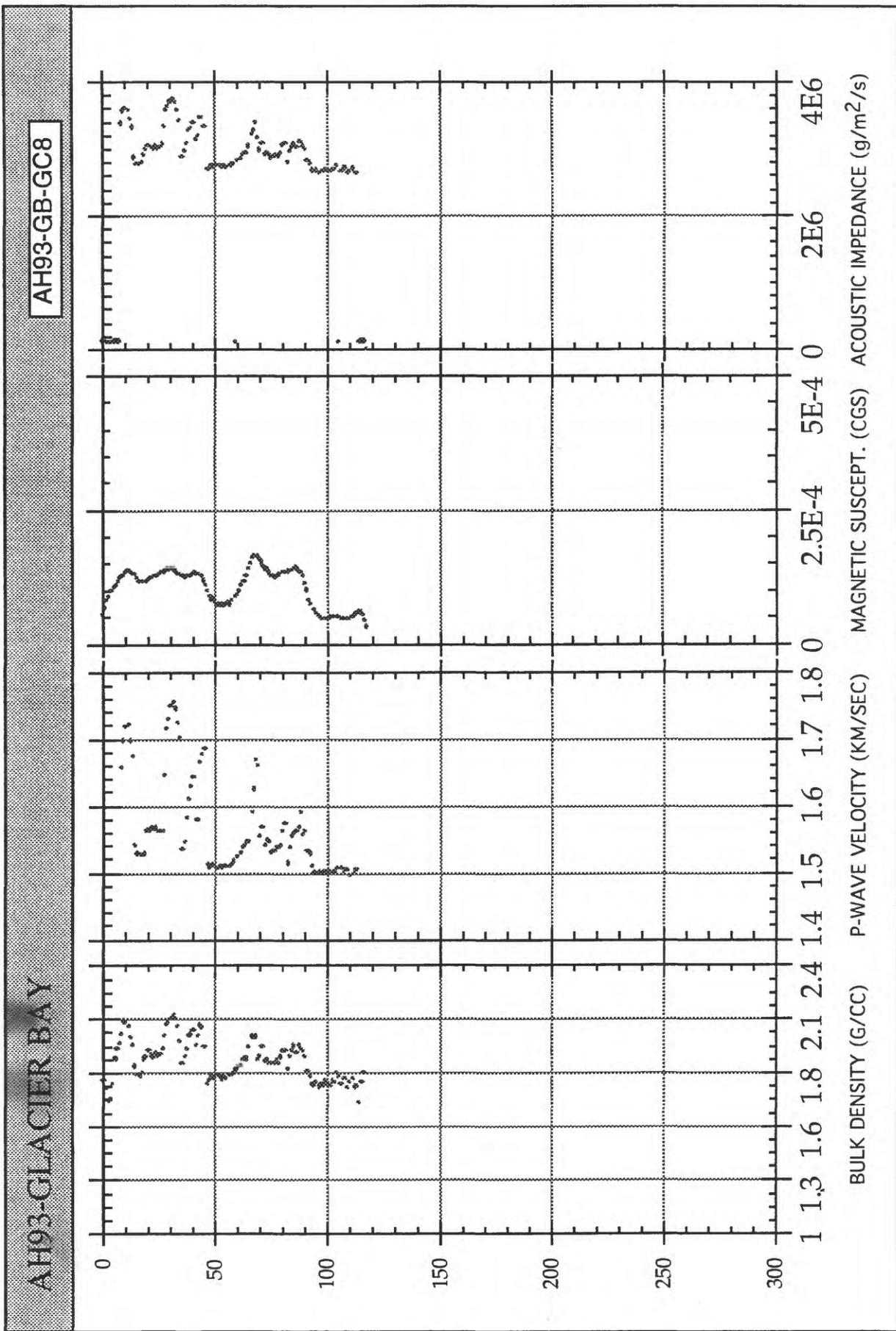


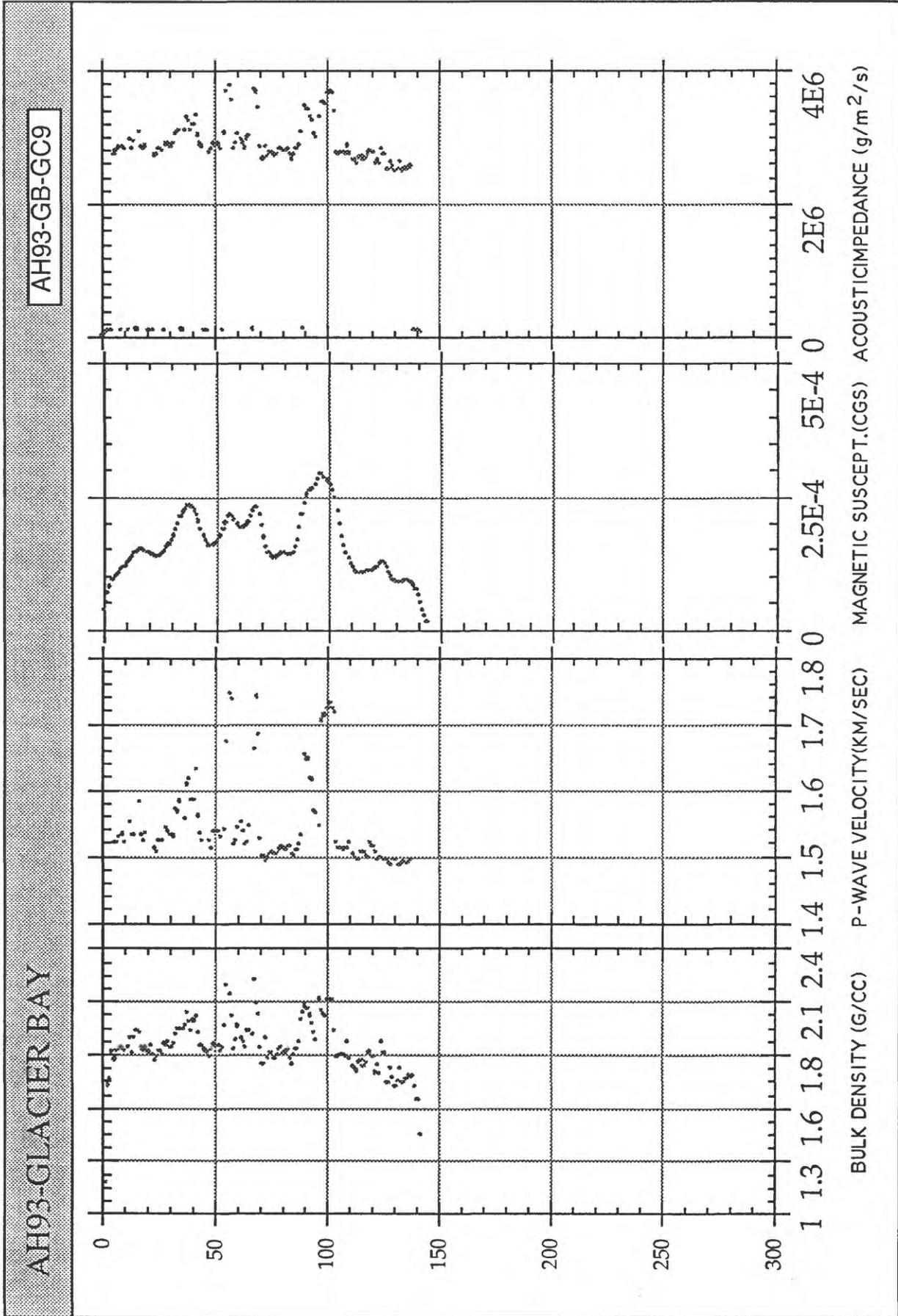


AH93-GLACIER BAY

AH93-GB-GC7

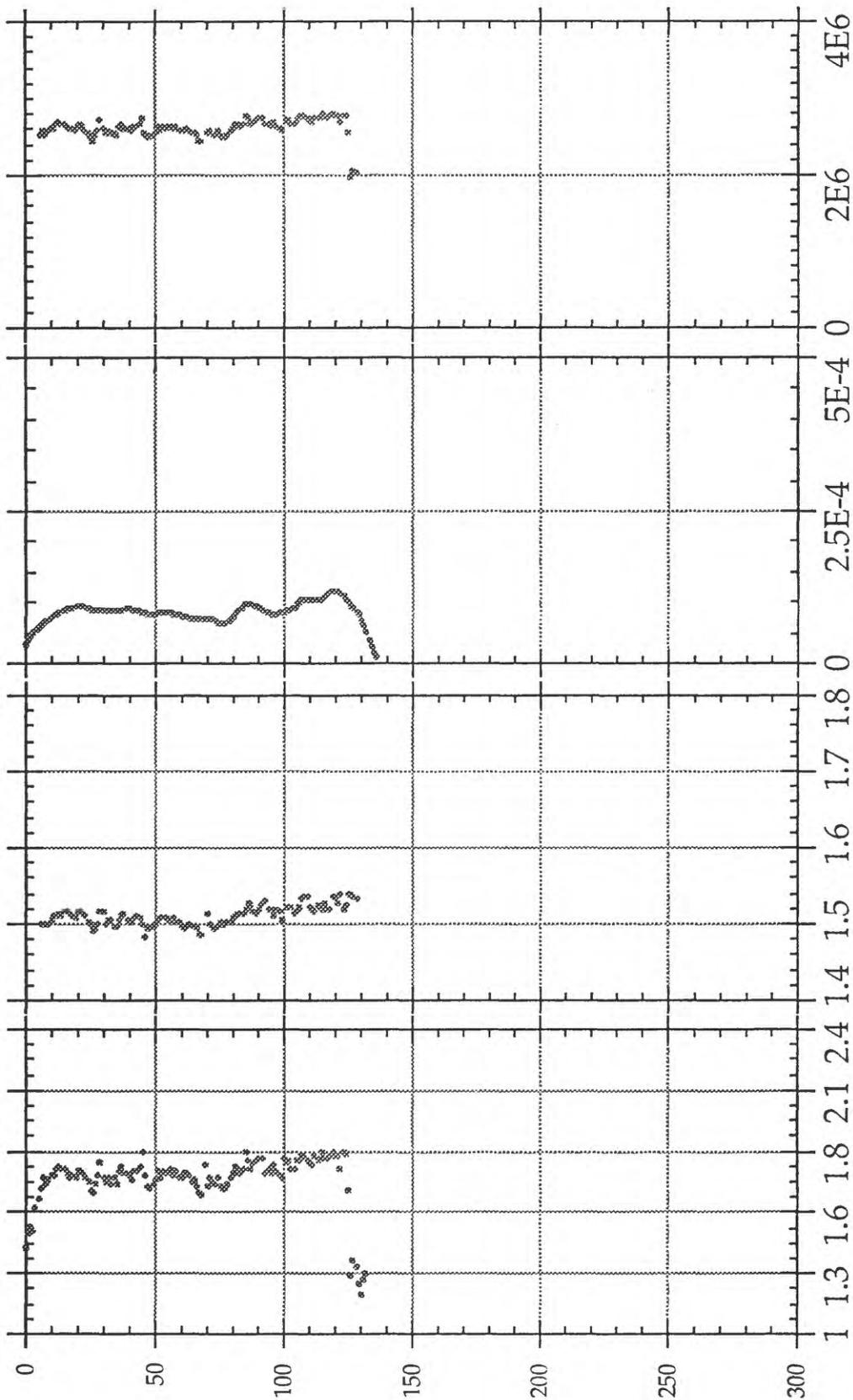




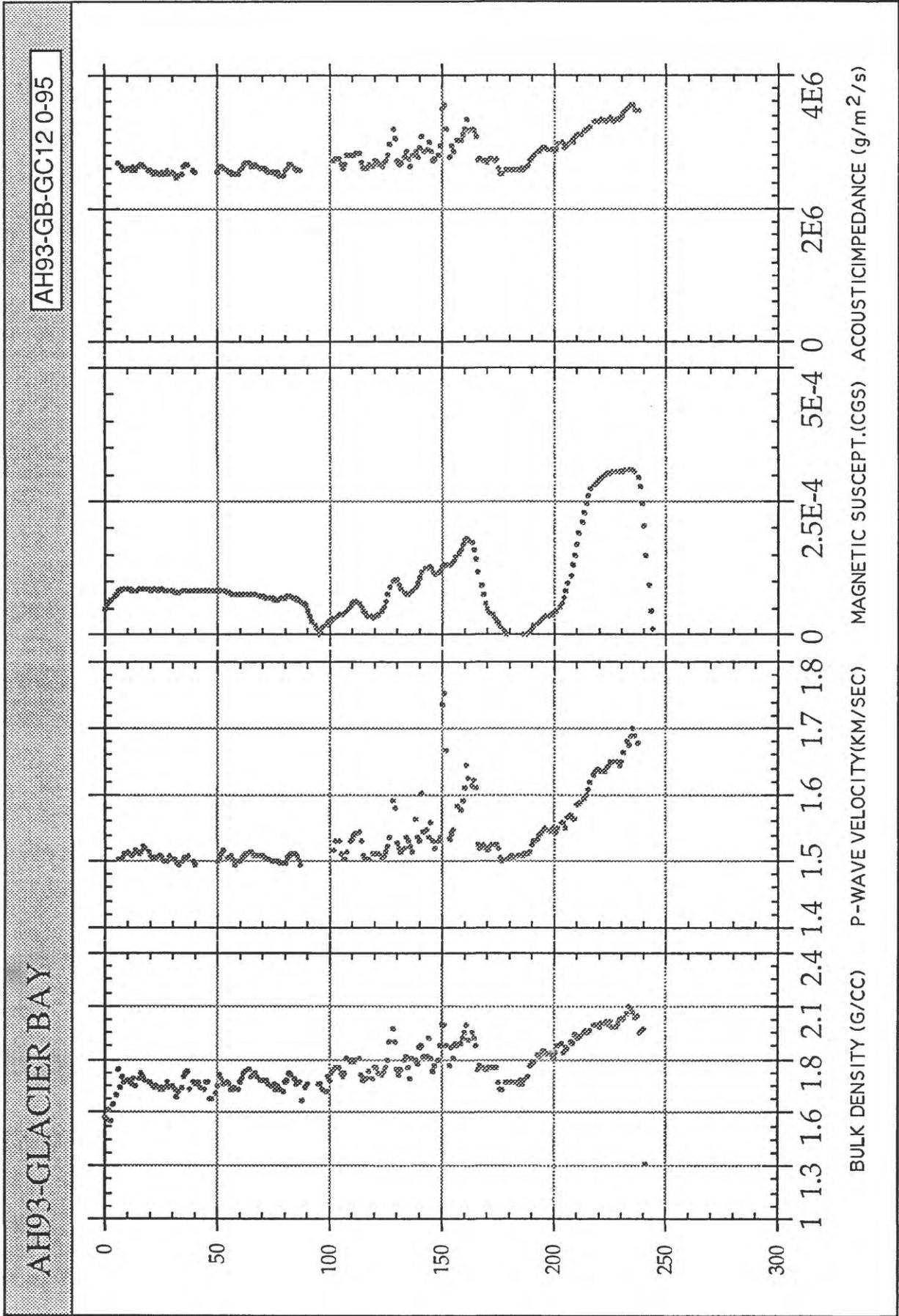


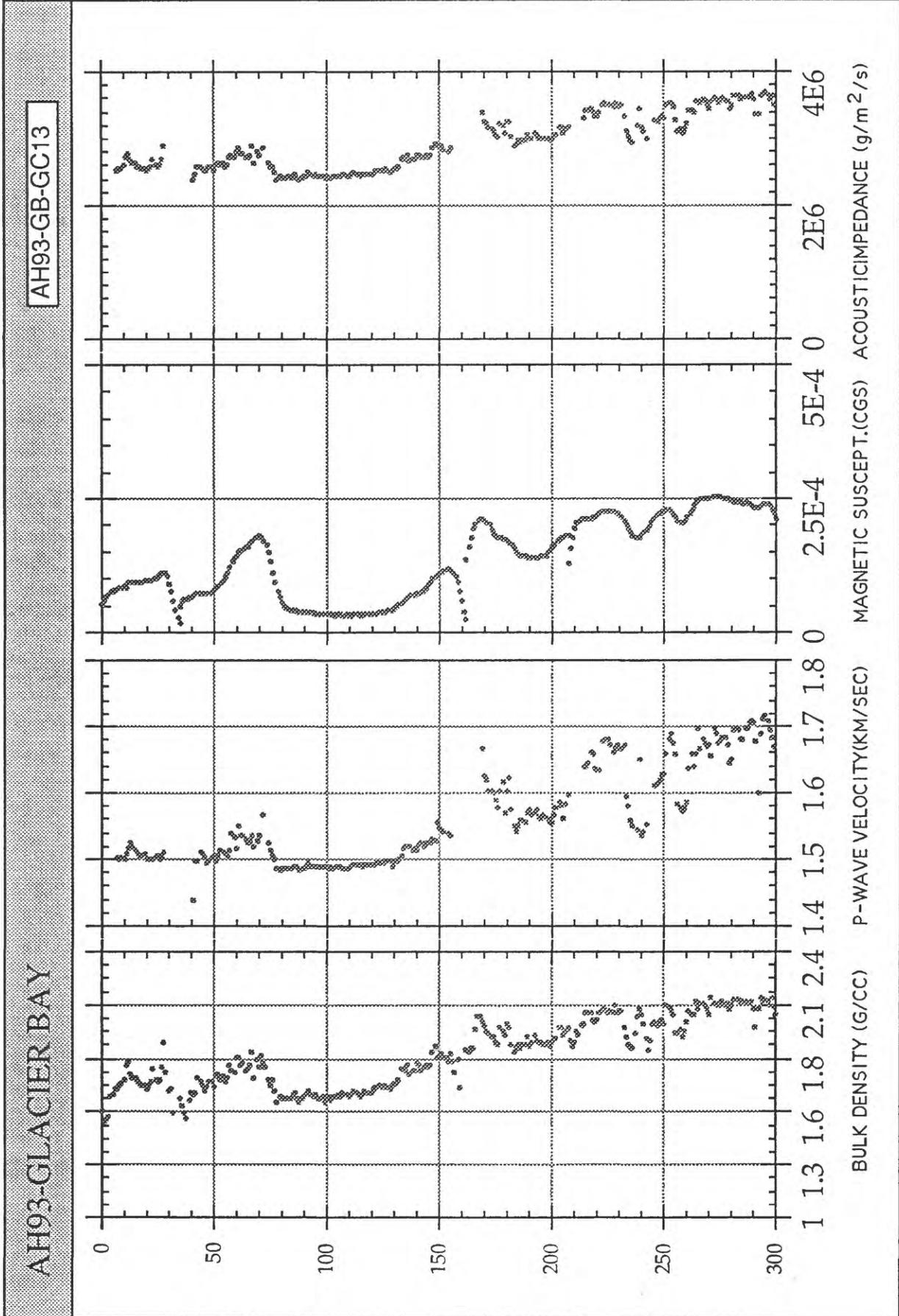
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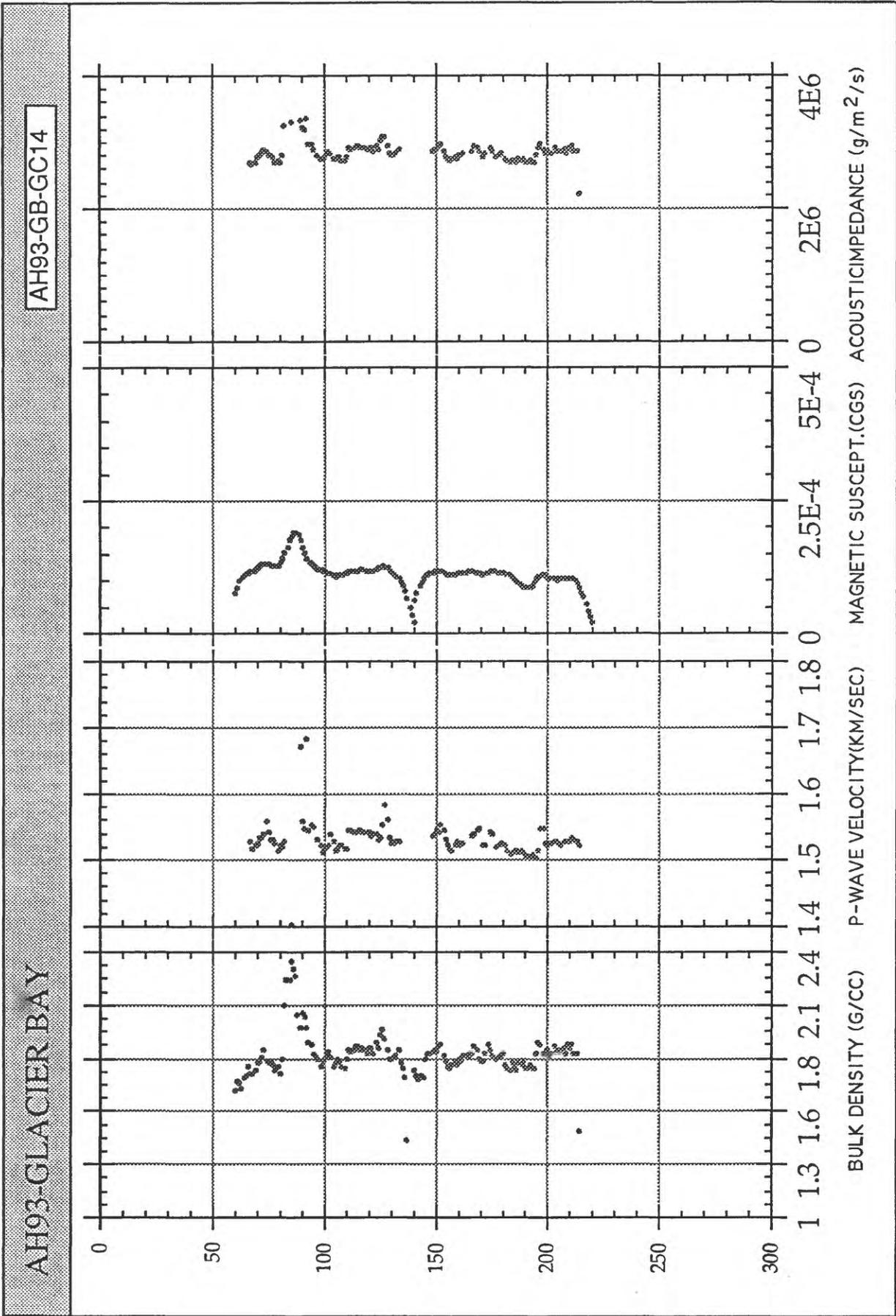
AH93-GB-GC11 0-136

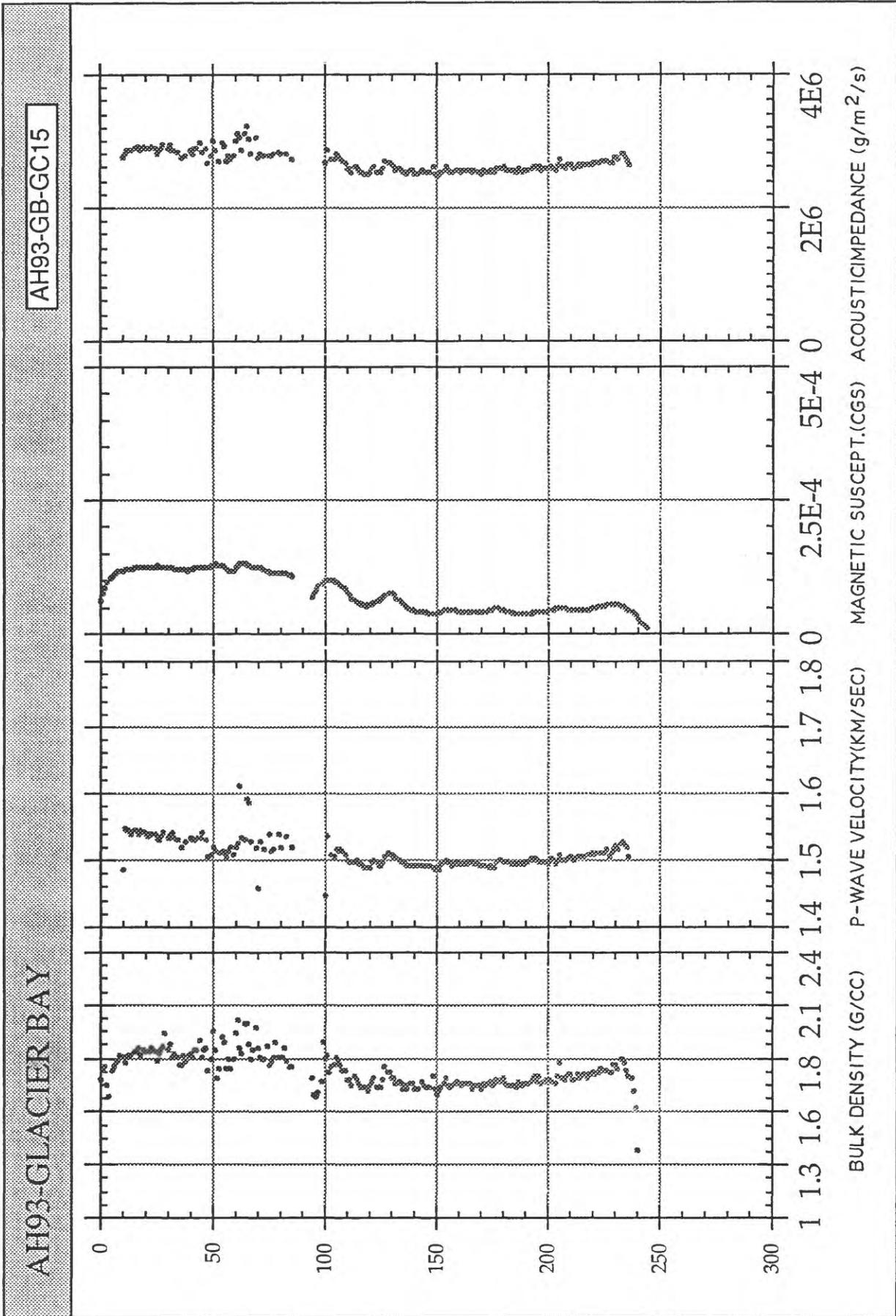


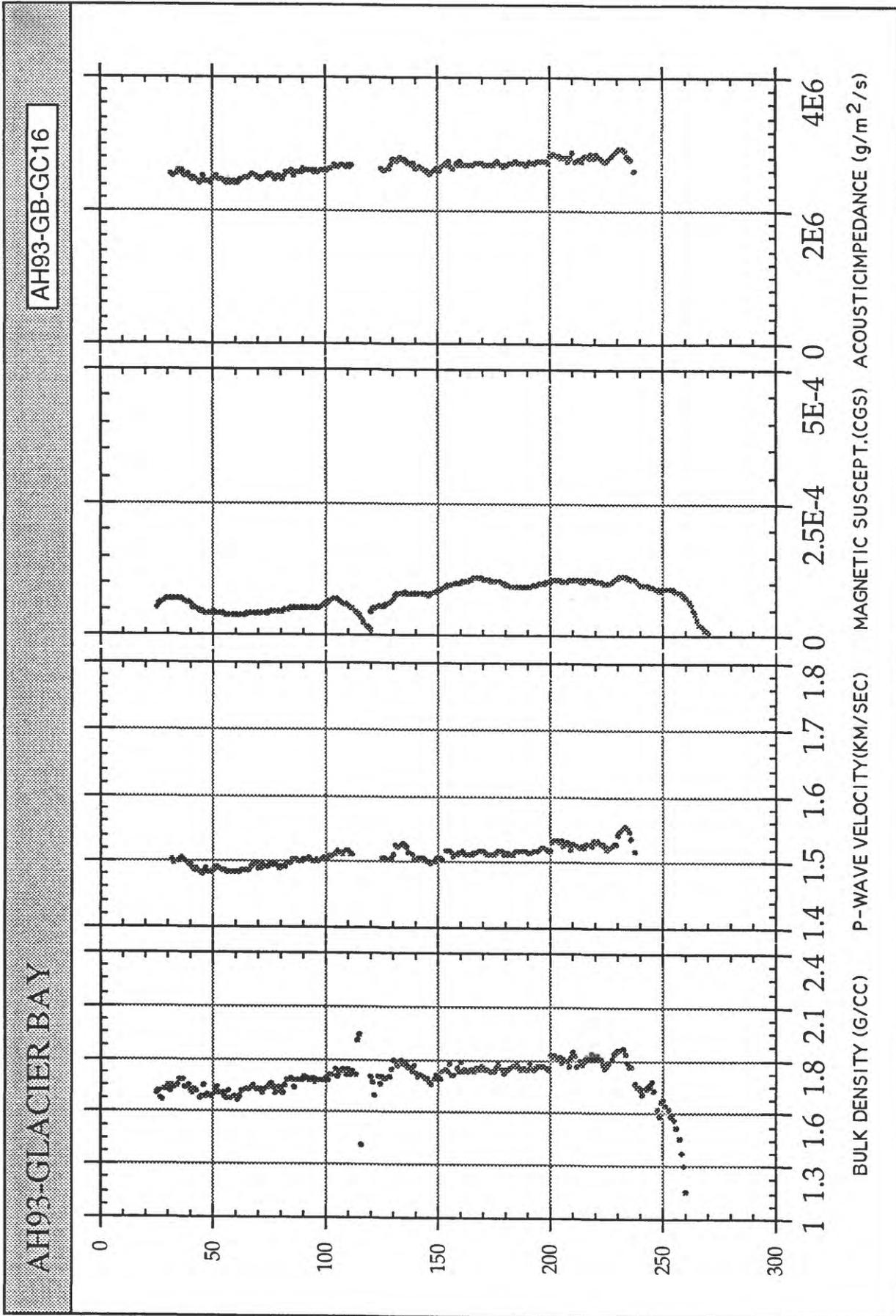
BULK DENSITY (G/CC) P-WAVE VELOCITY(KM/SEC) MAGNETIC SUSCEPT.(CGS) ACOUSTICIMPEDANCE (g/m²/s)

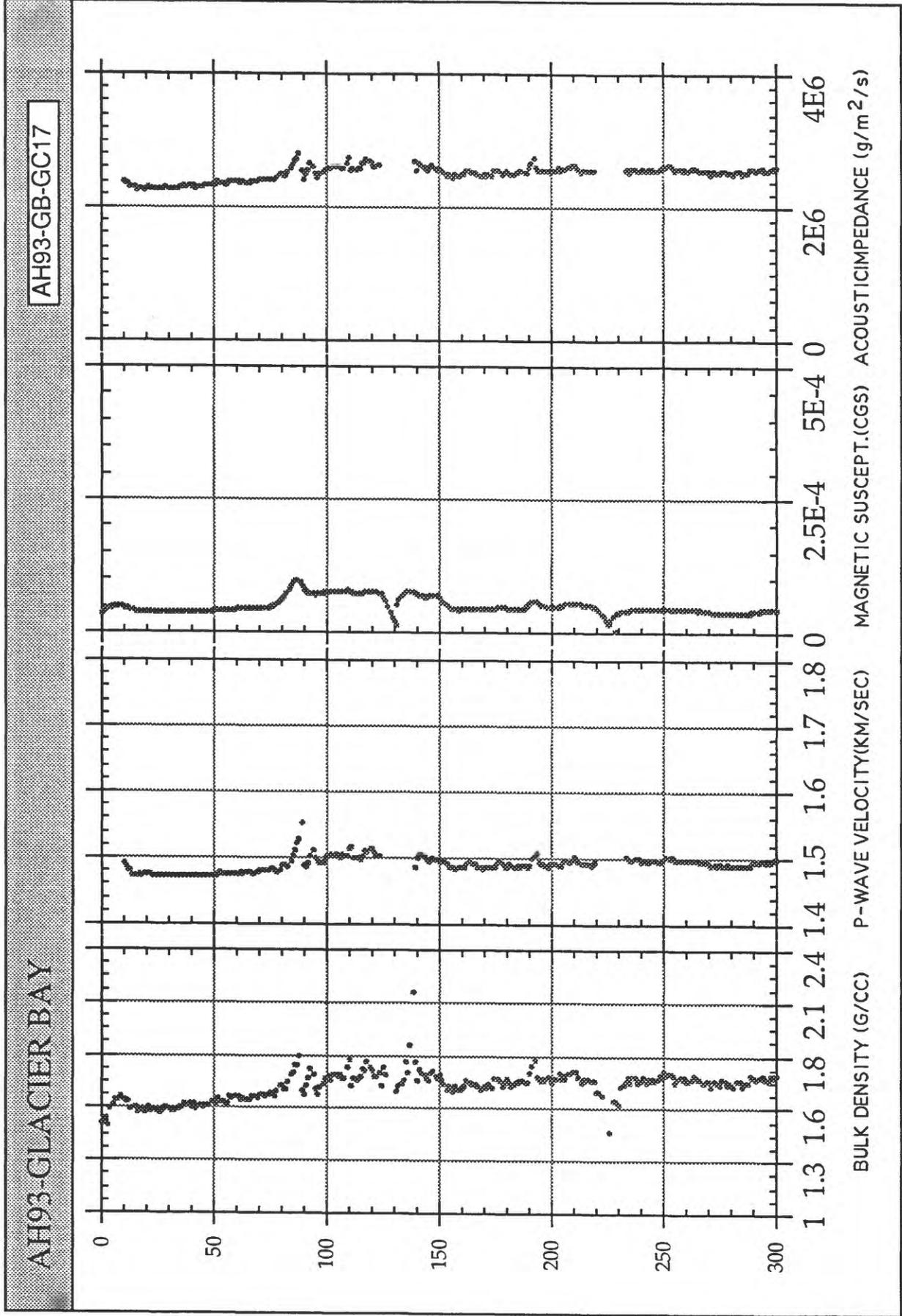


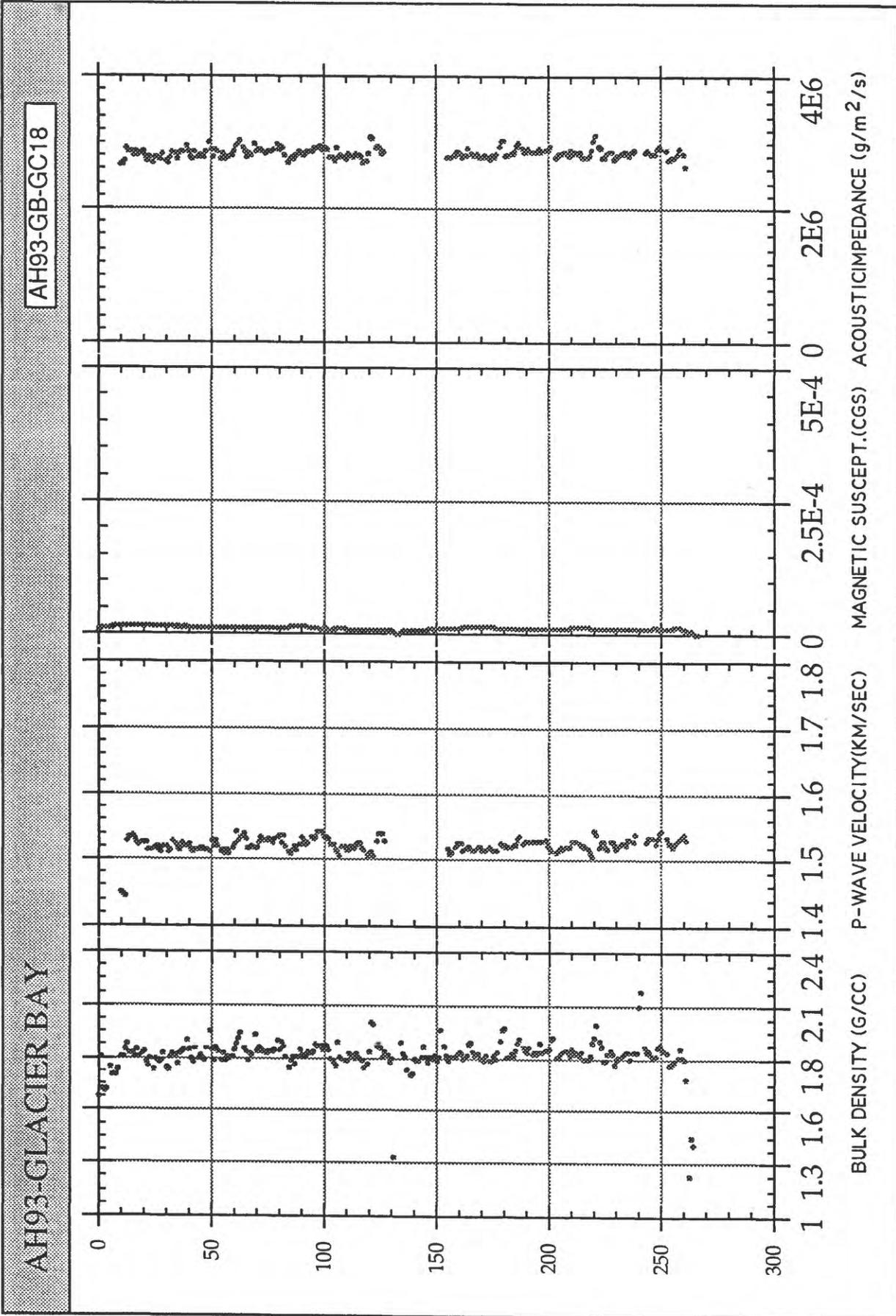






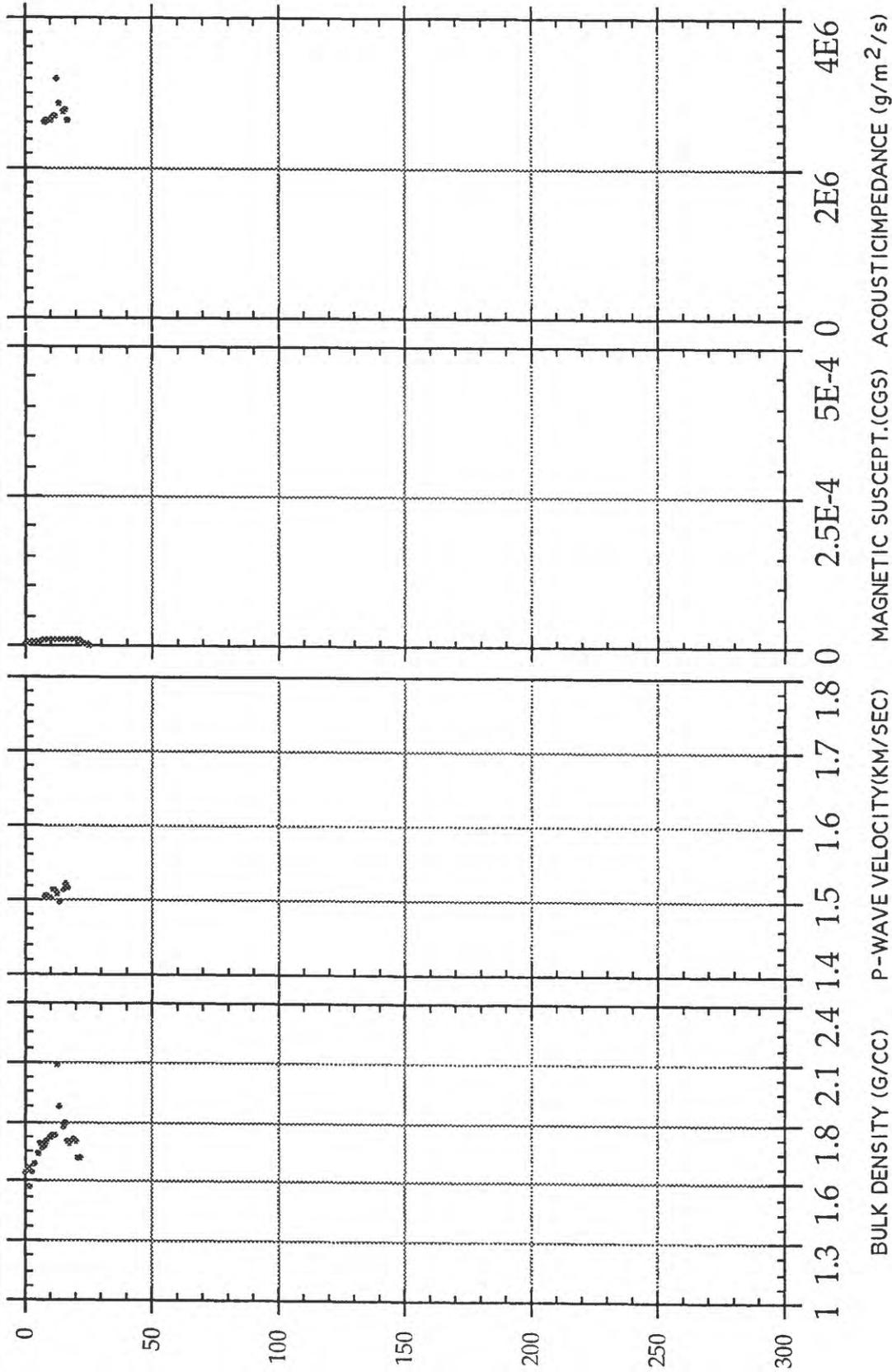


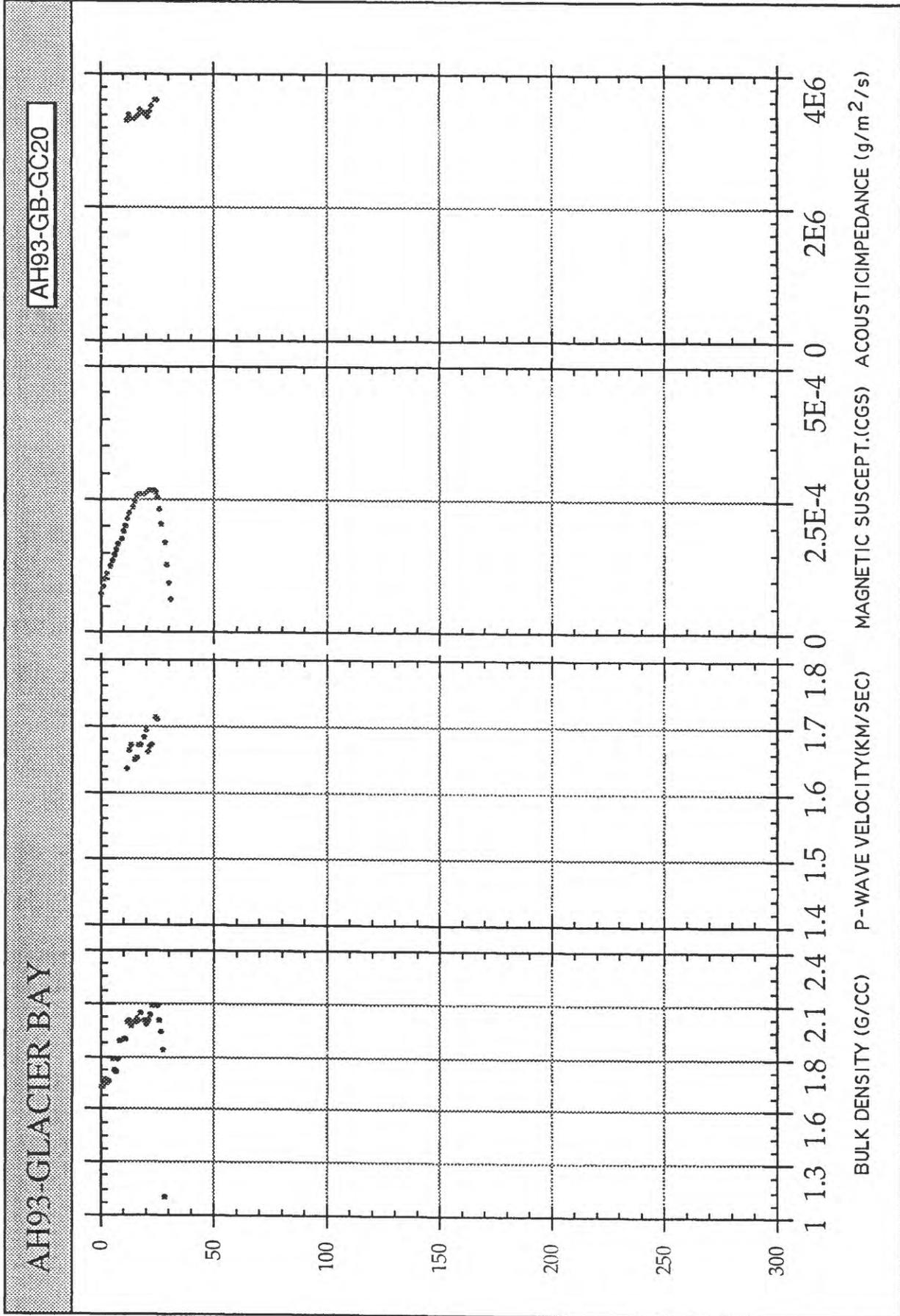


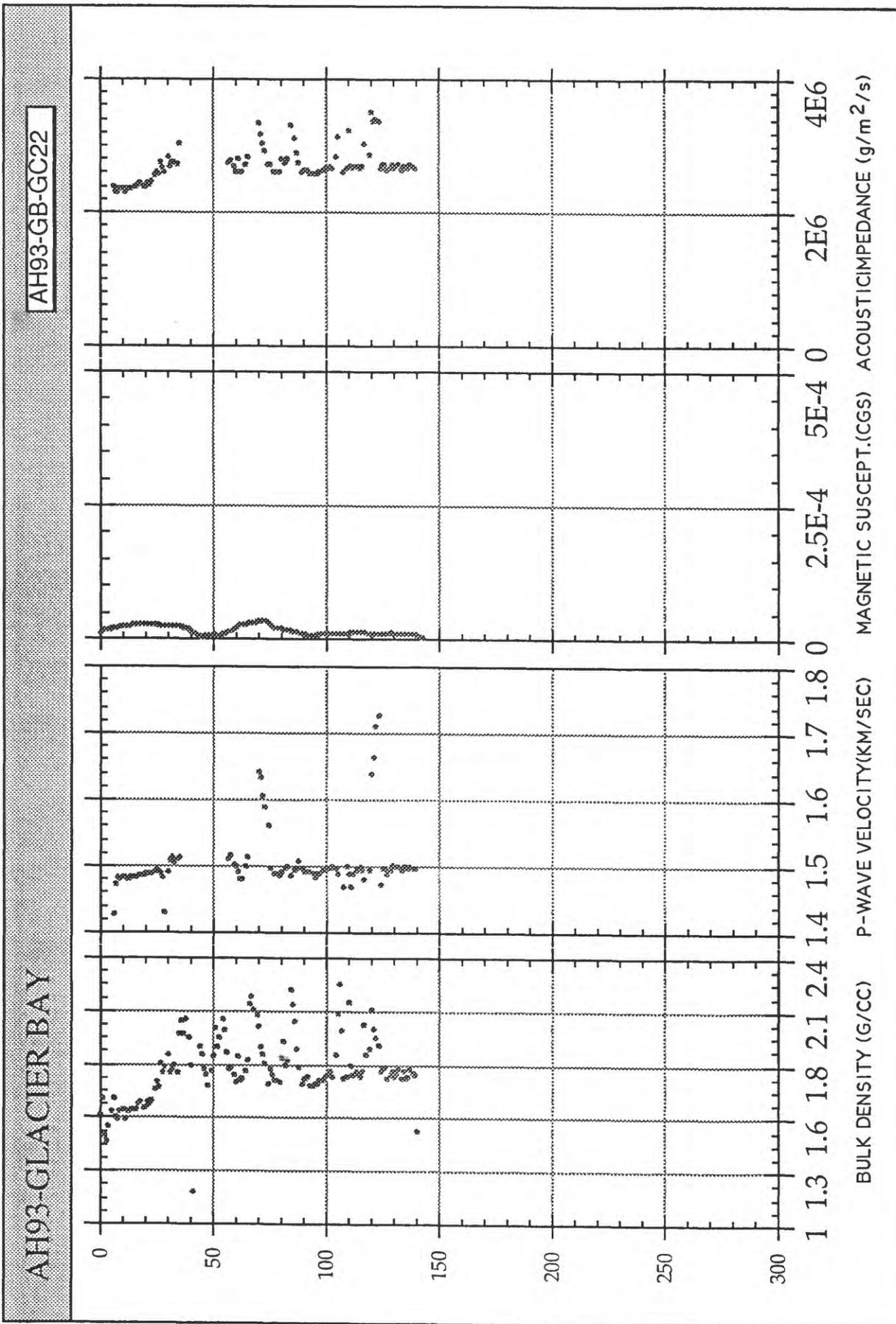


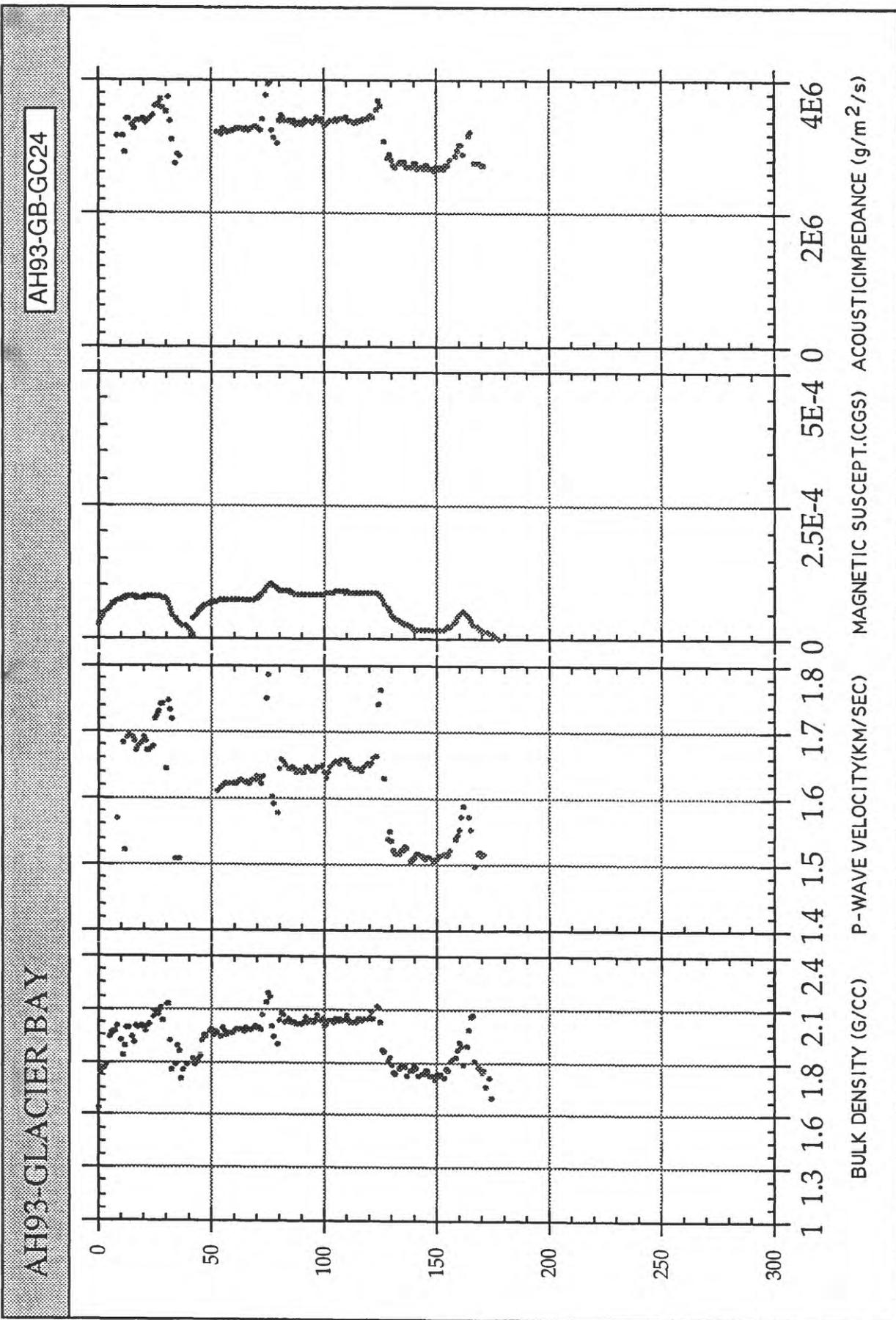
AH93-GLACIER BAY

AH93-GB-GC19



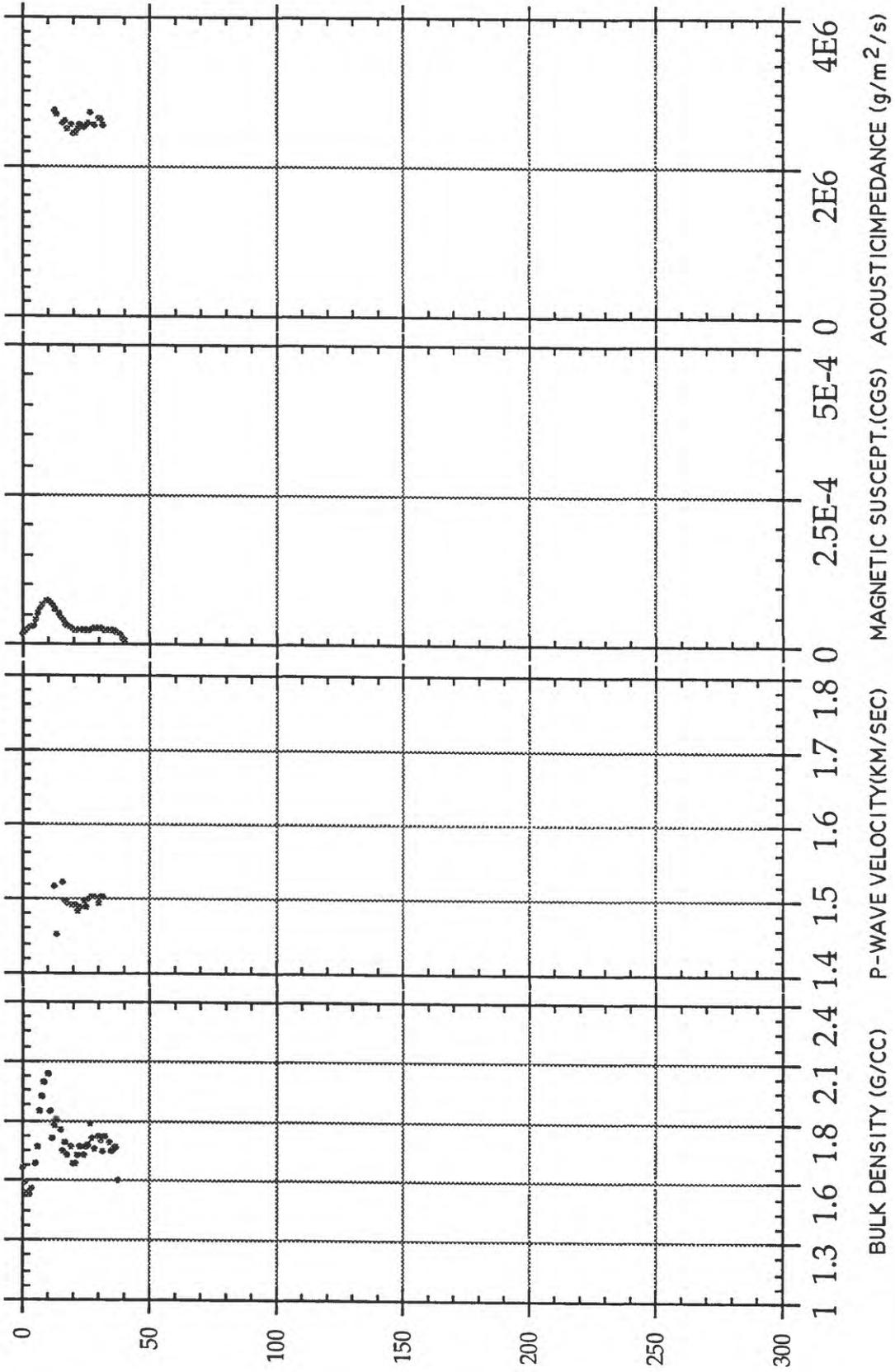


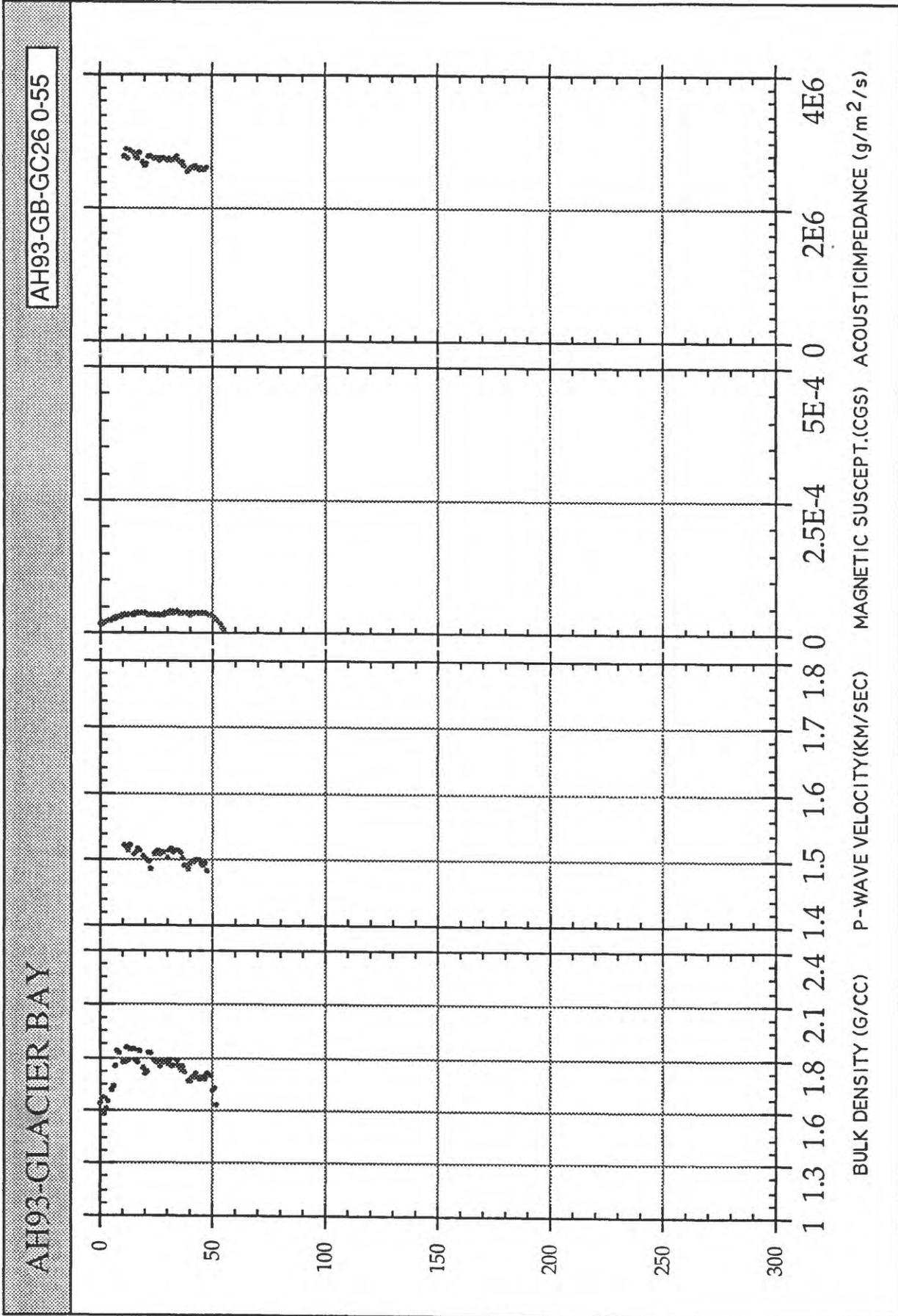


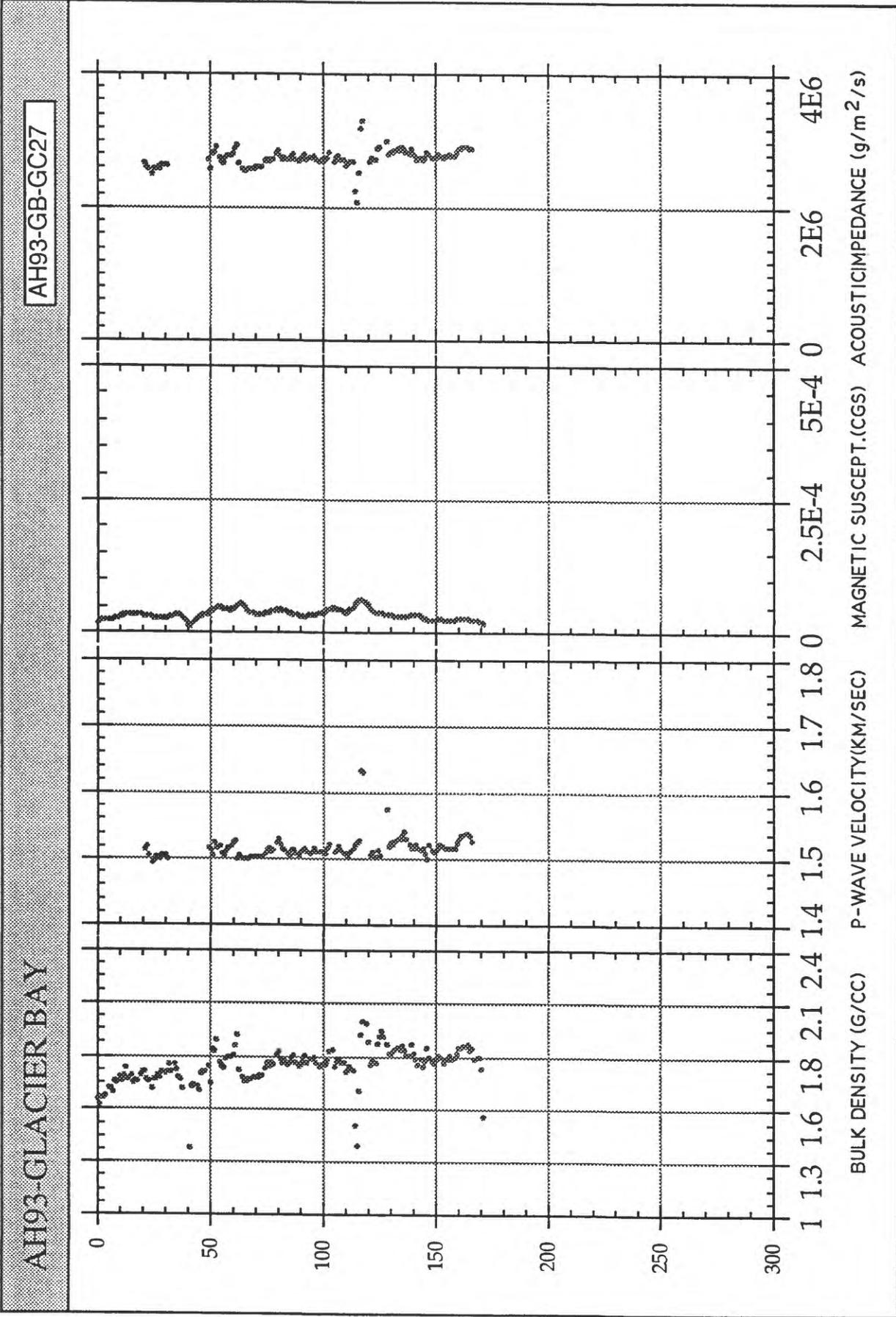


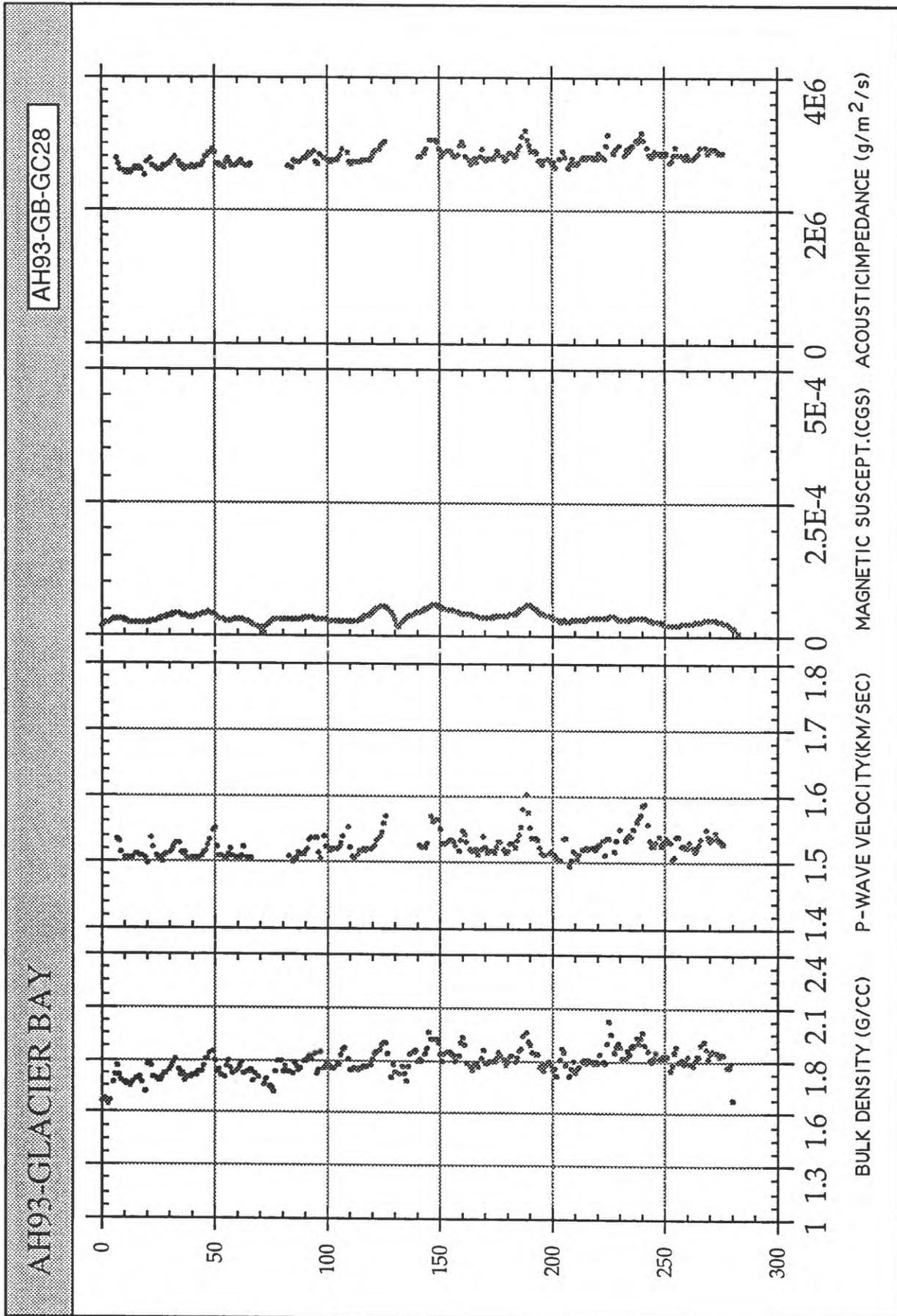
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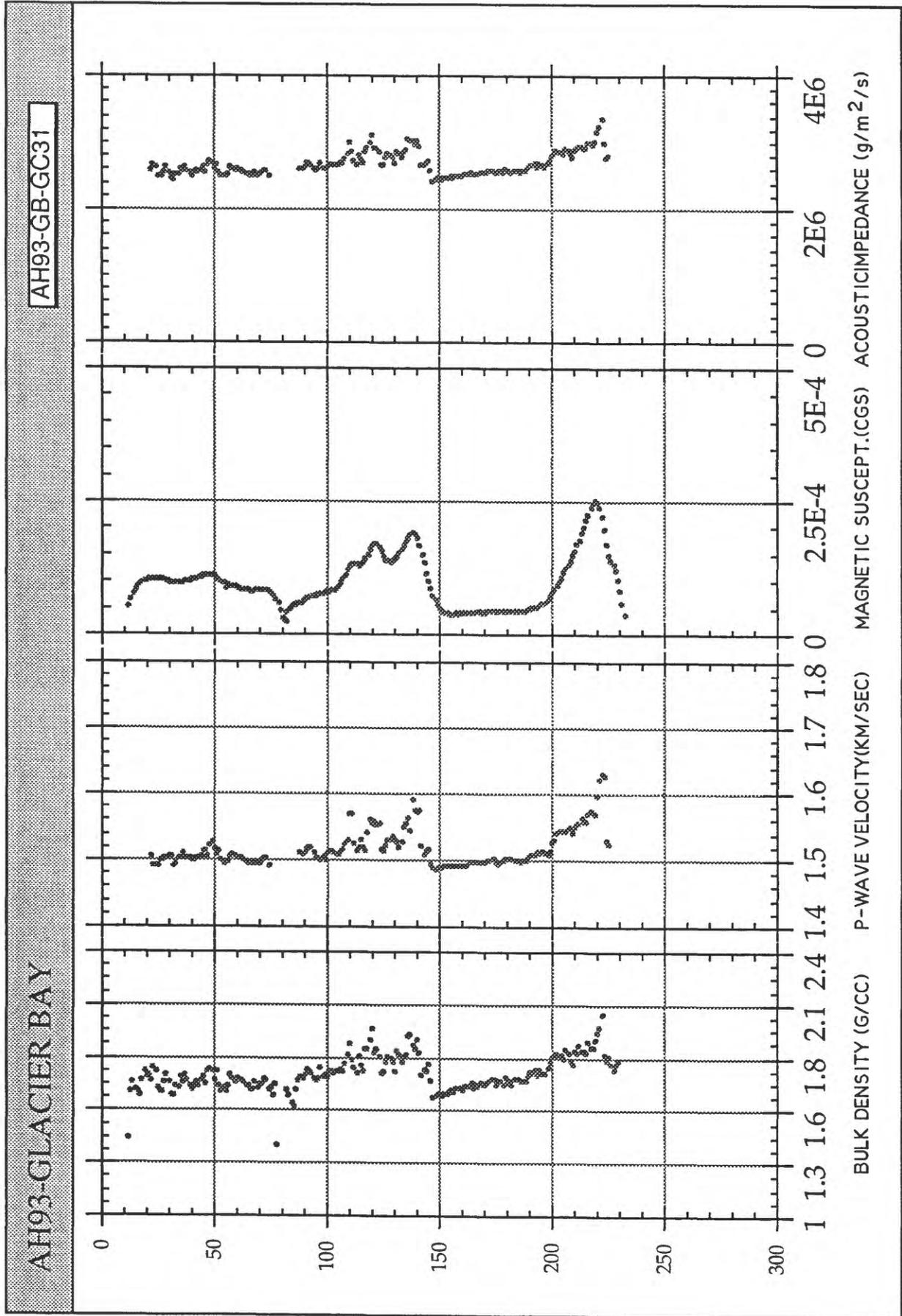
AH93-GB-GC25 0-40

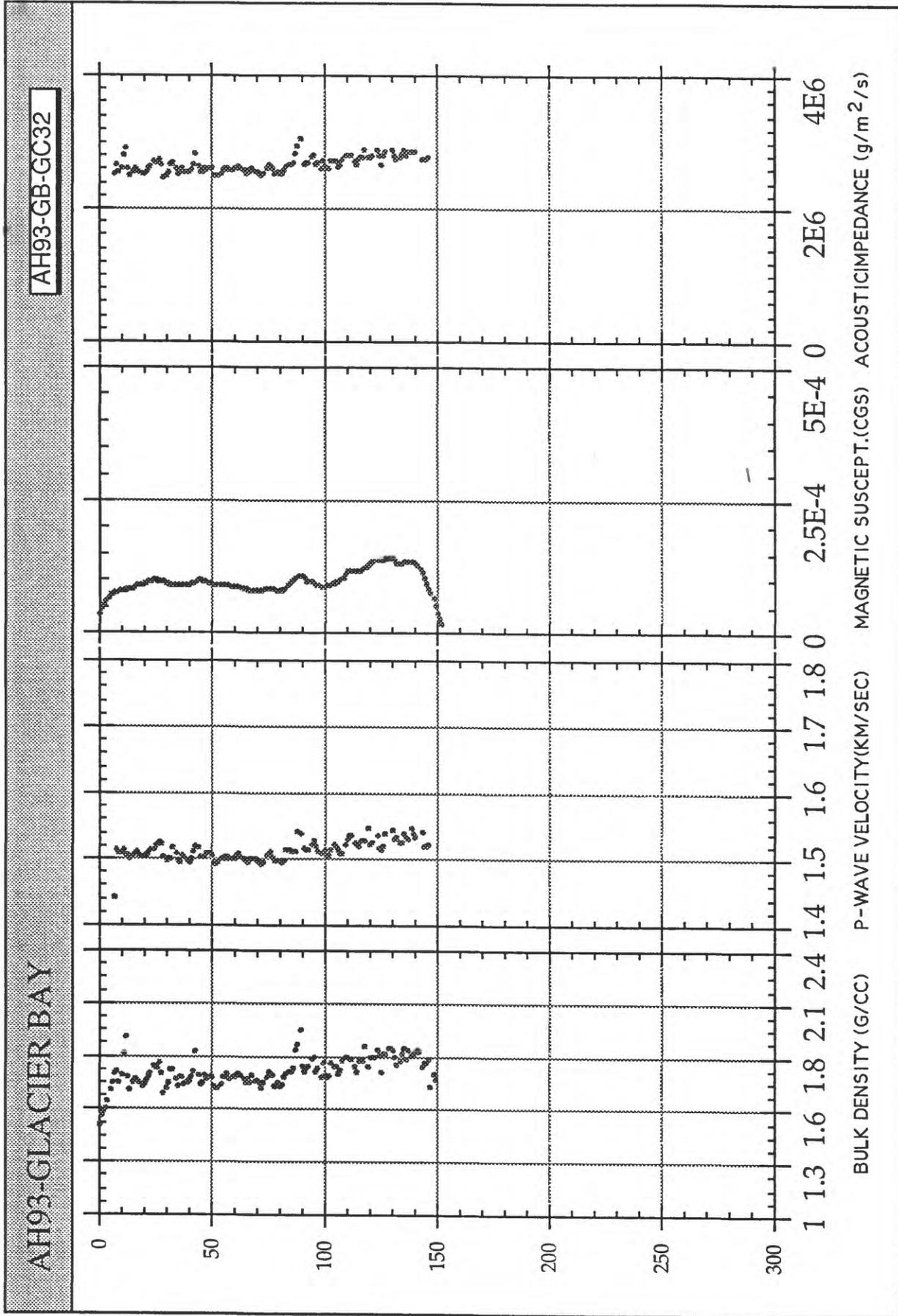






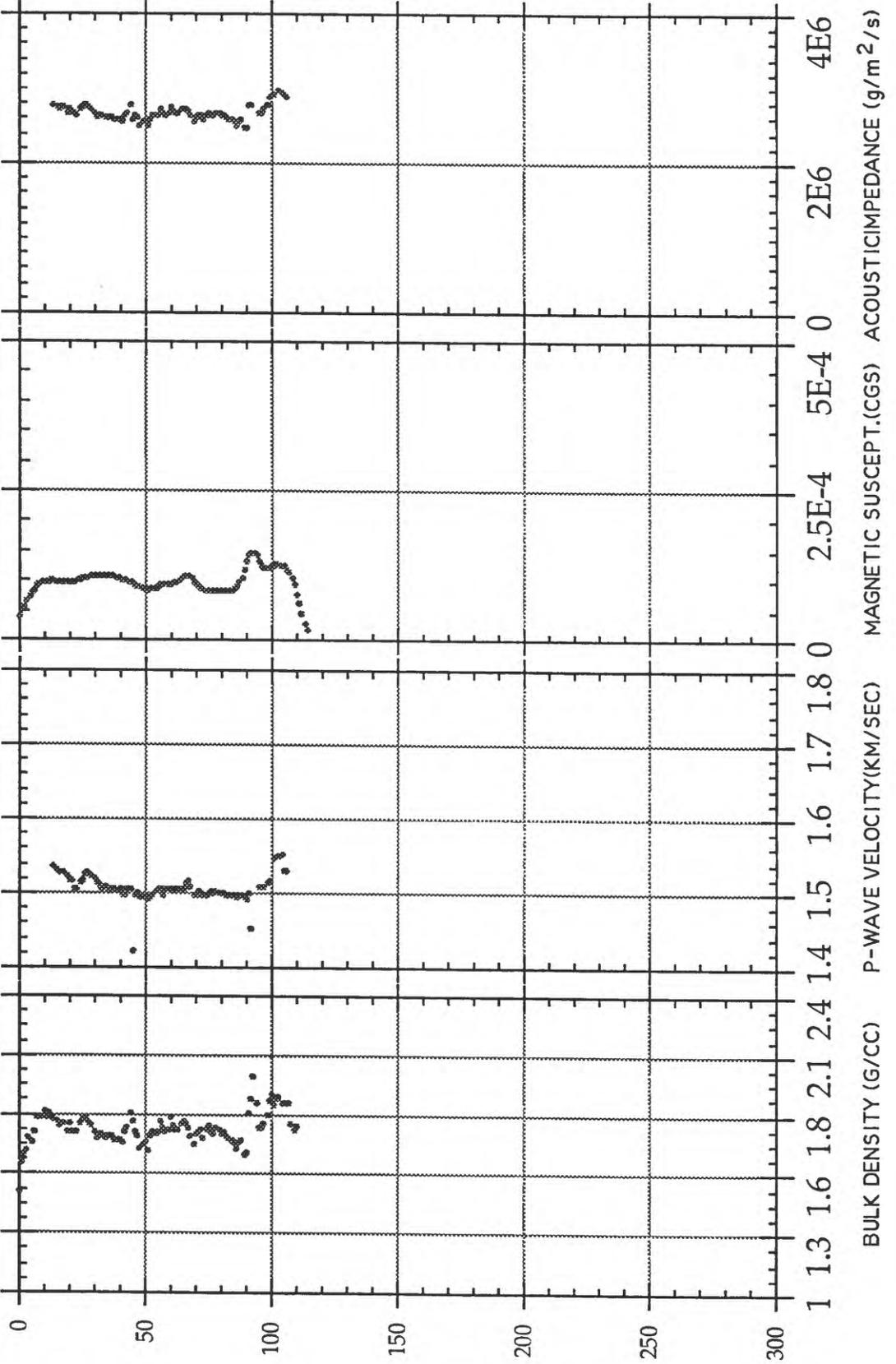






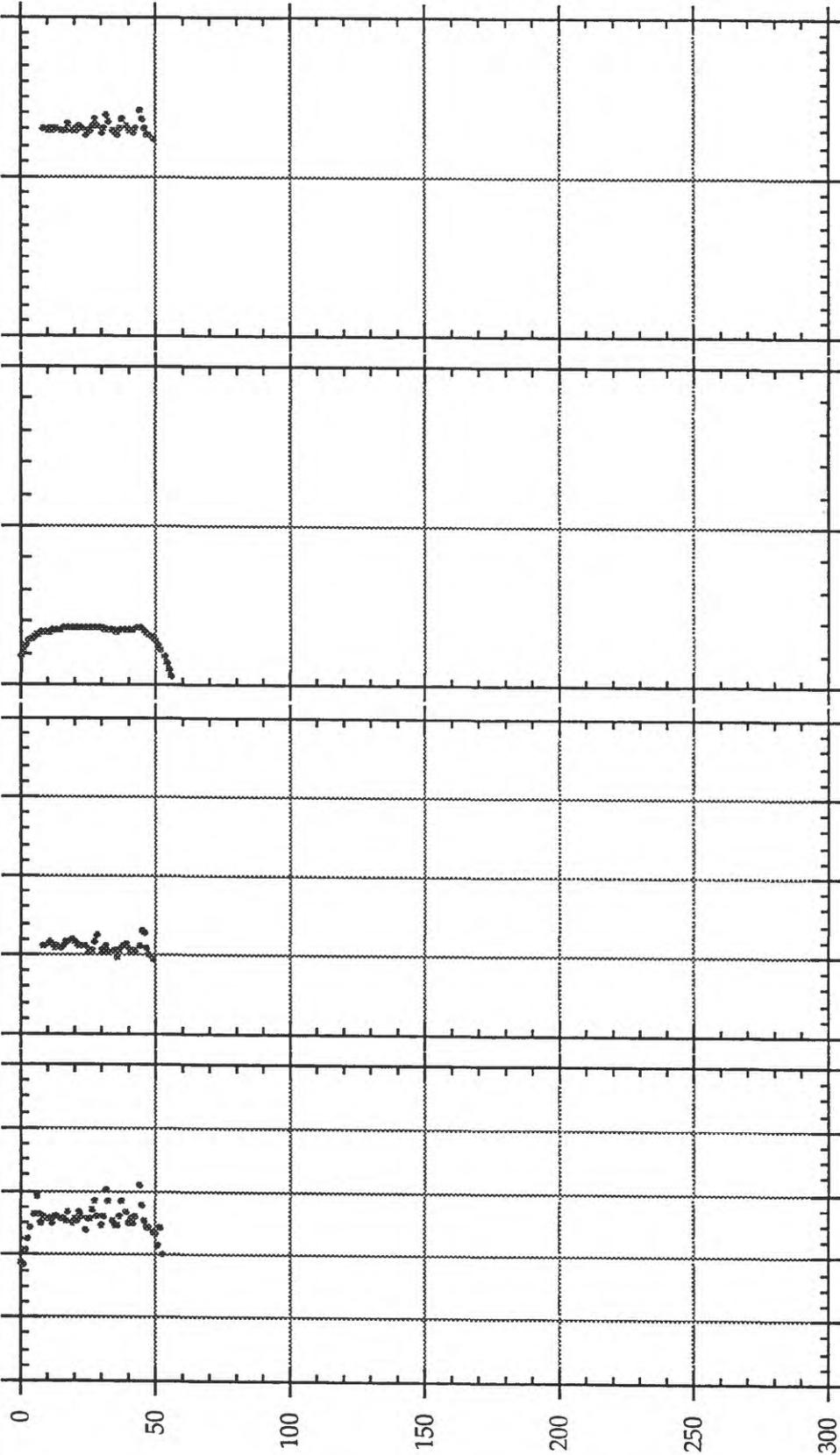
AH93-GB-GC33

AH93-GLACIER BAY

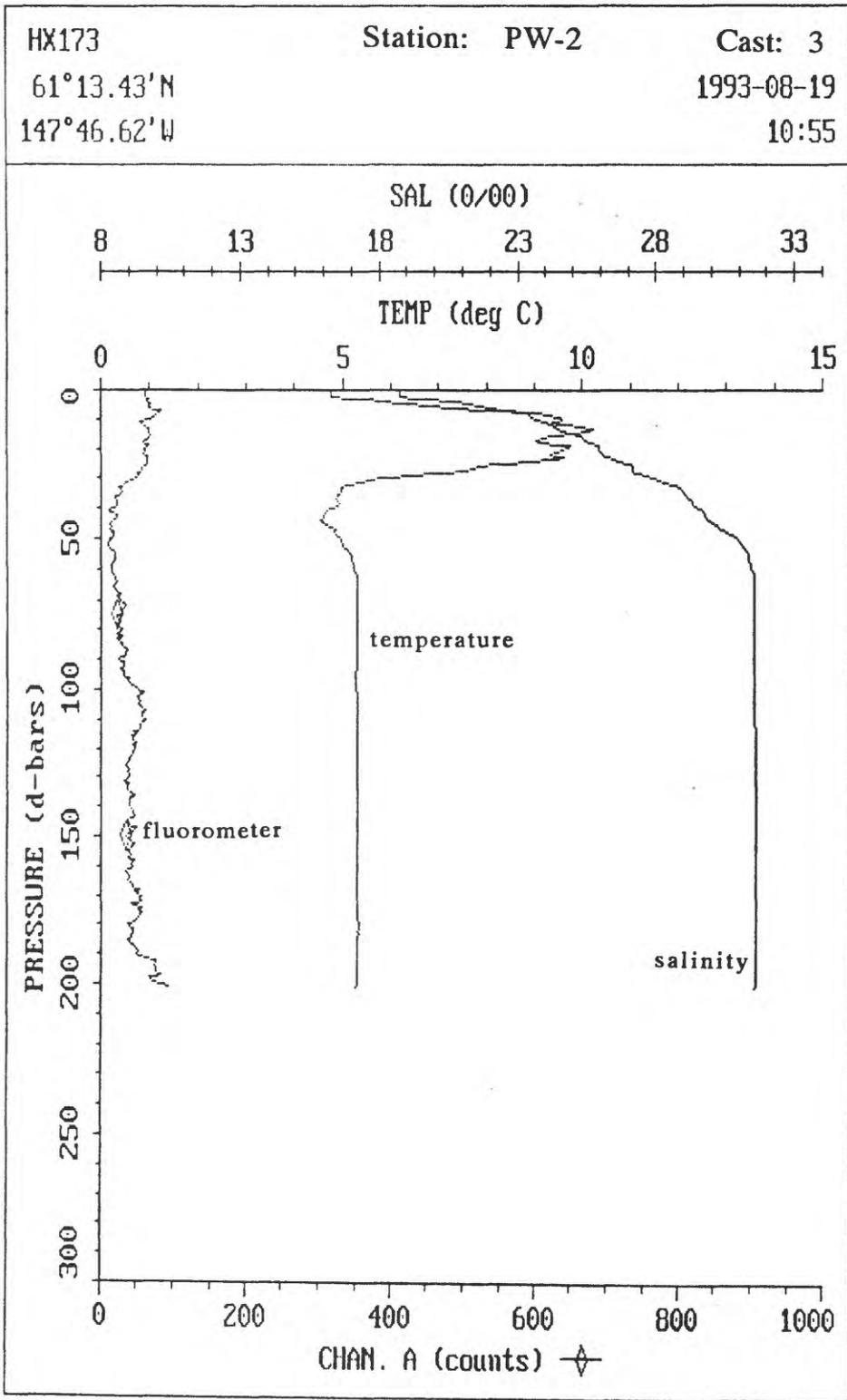


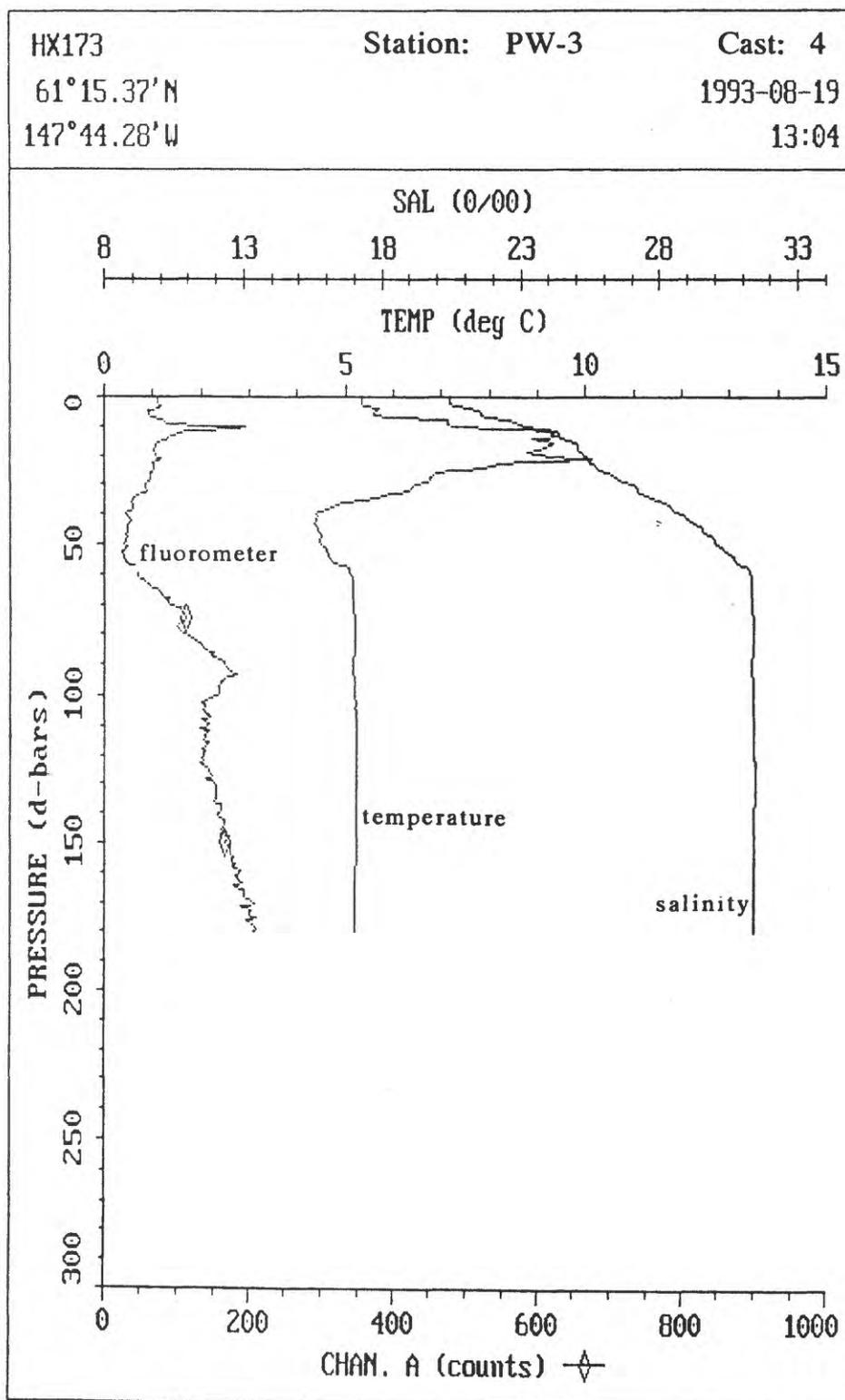
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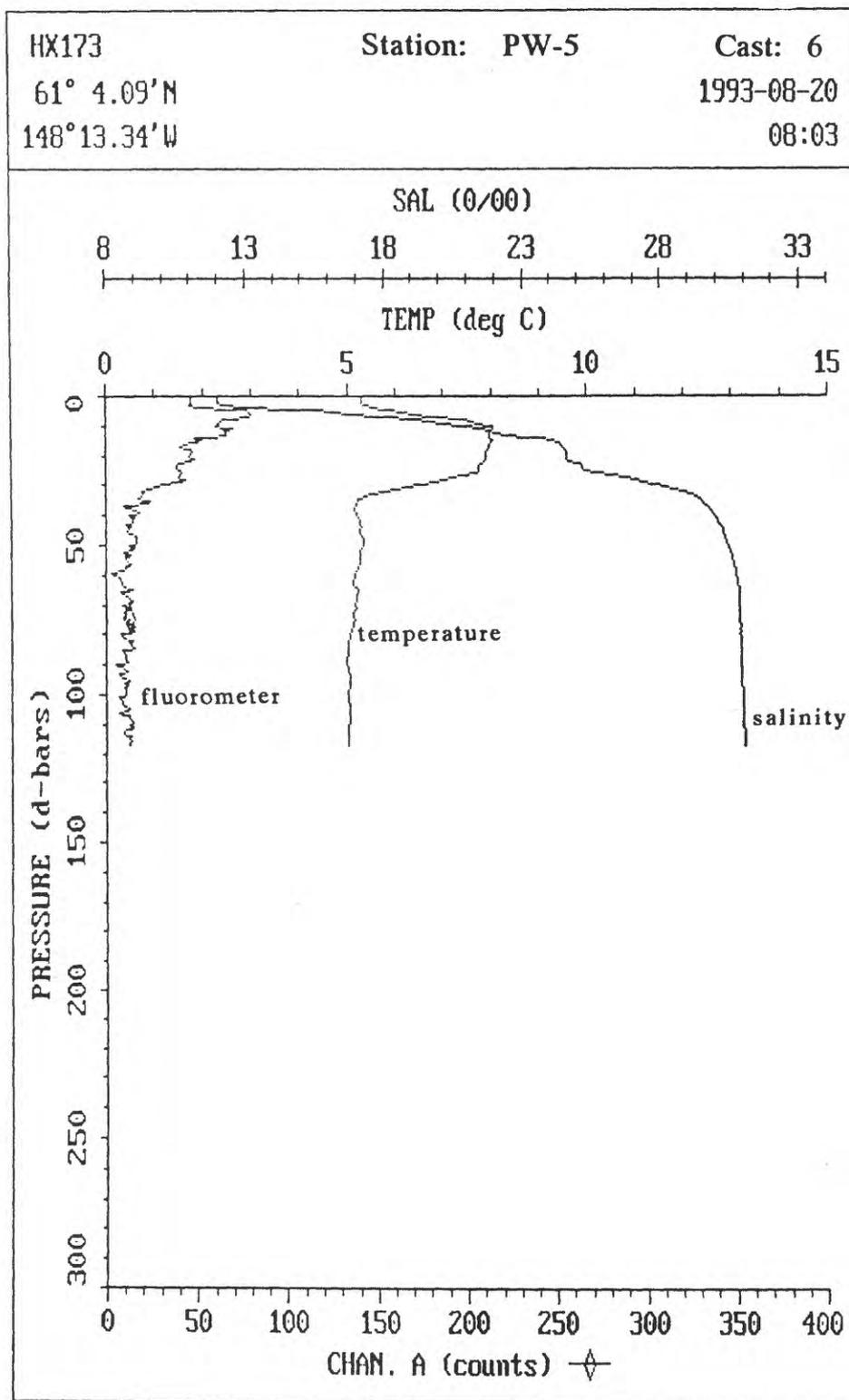
AH93-GLACIER BAY

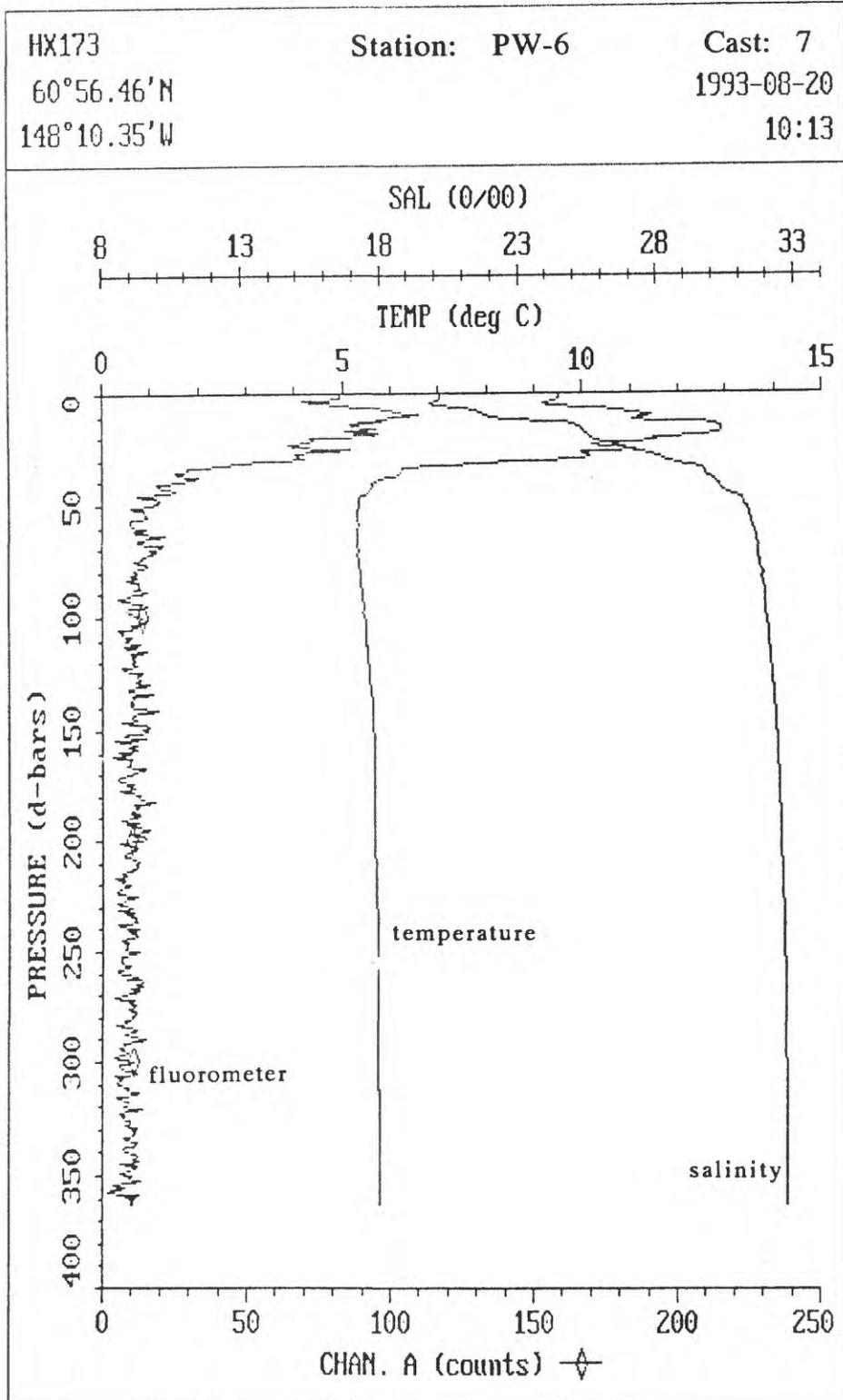


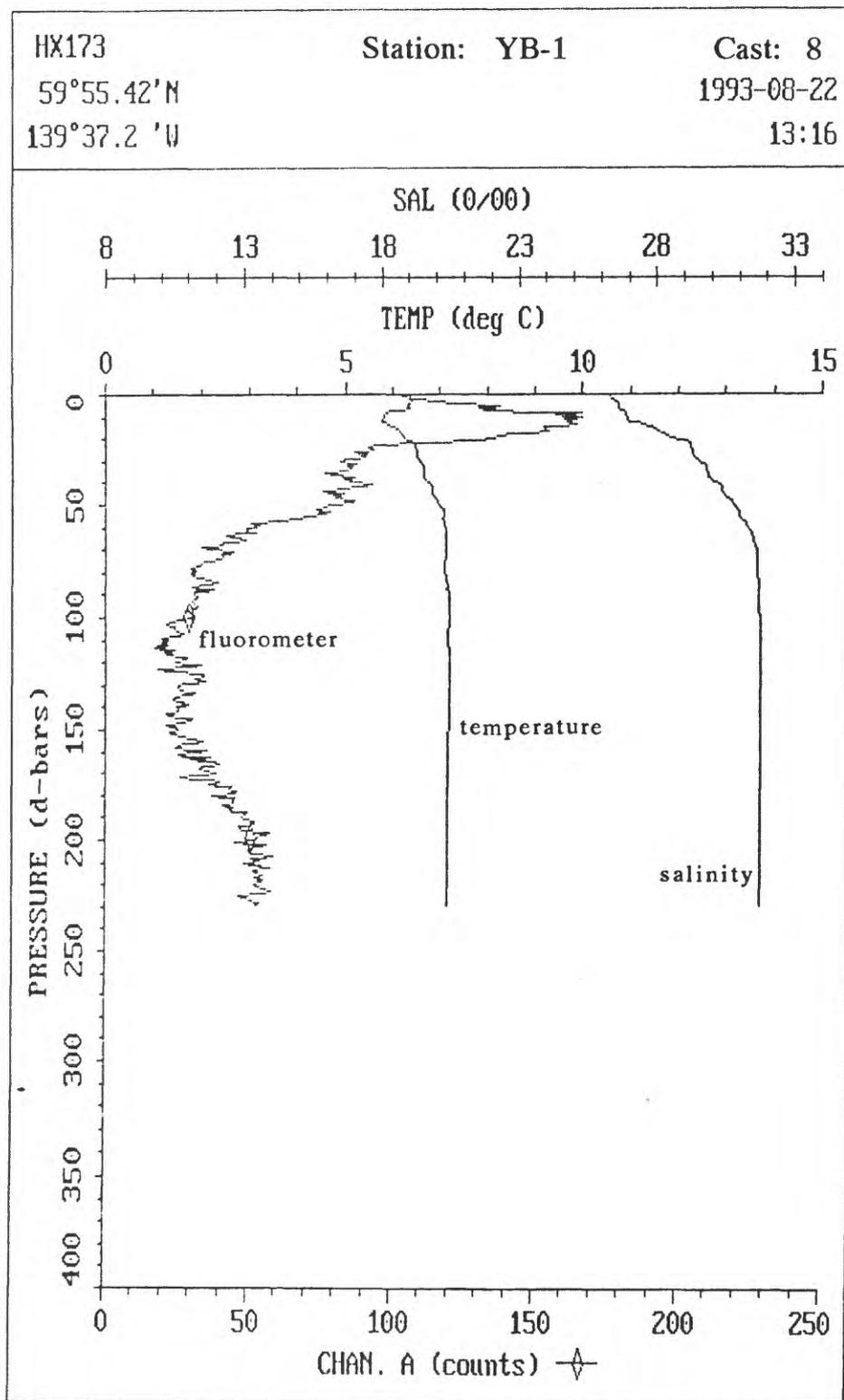
APPENDIX 3: CTD PROFILES

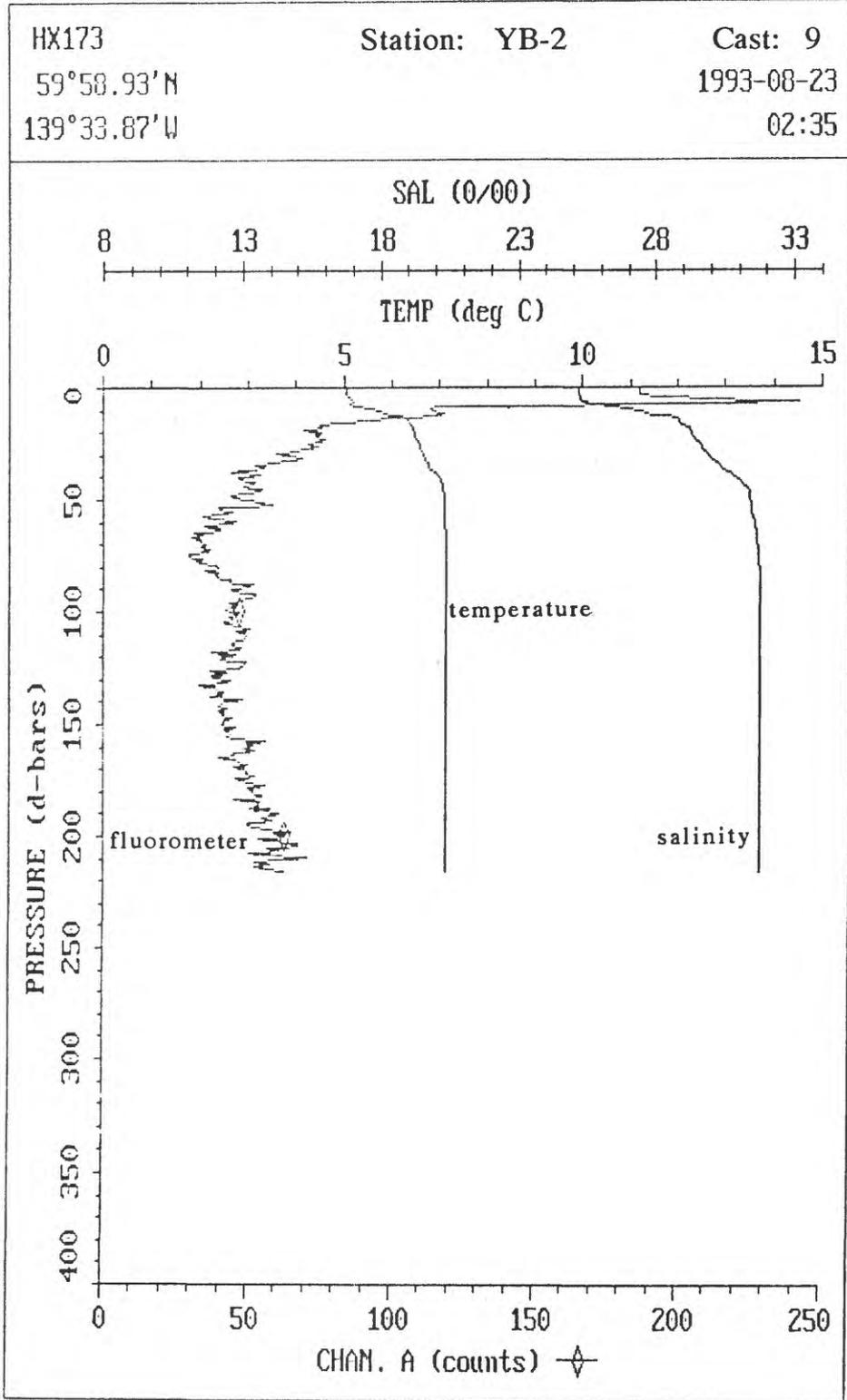


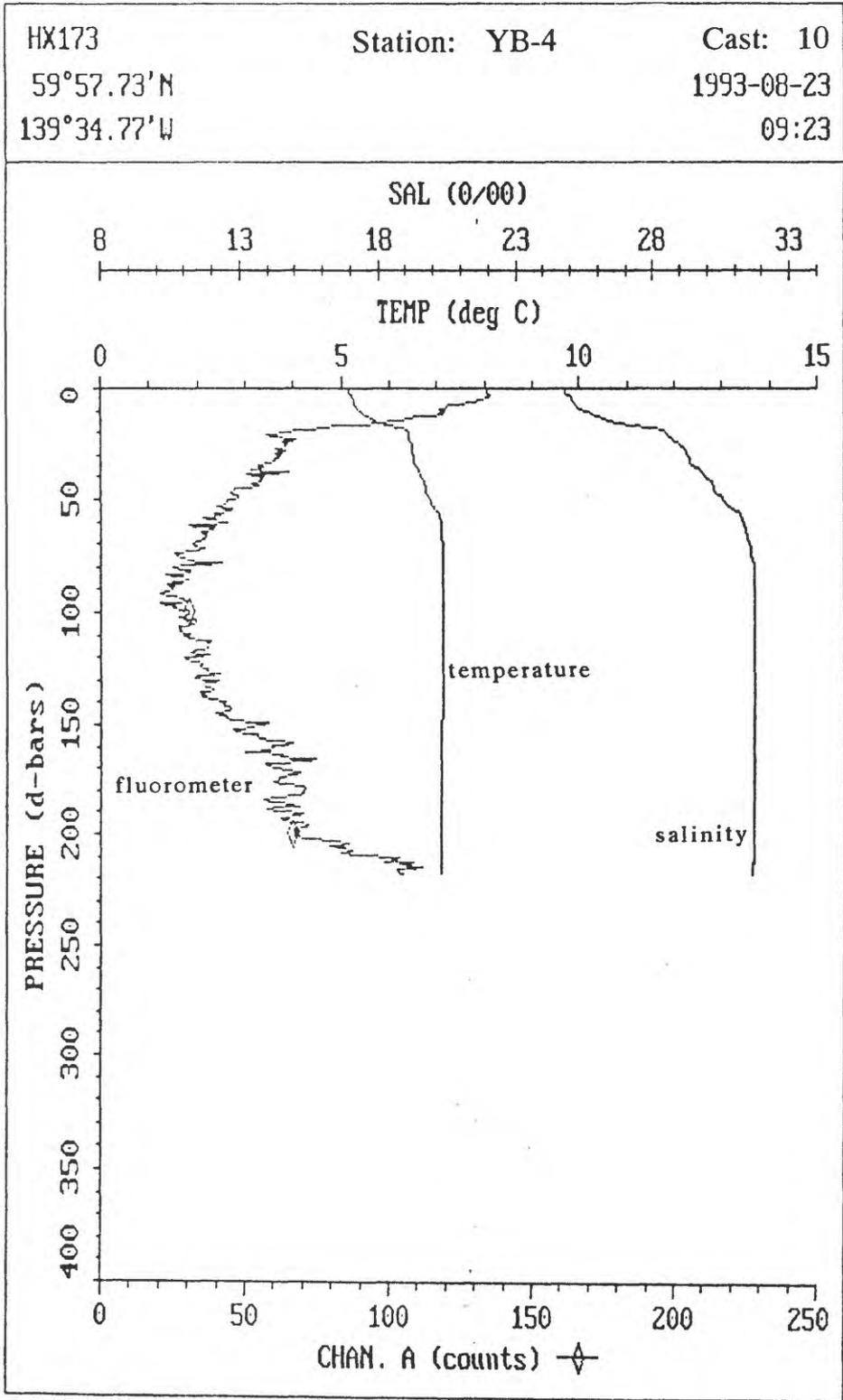


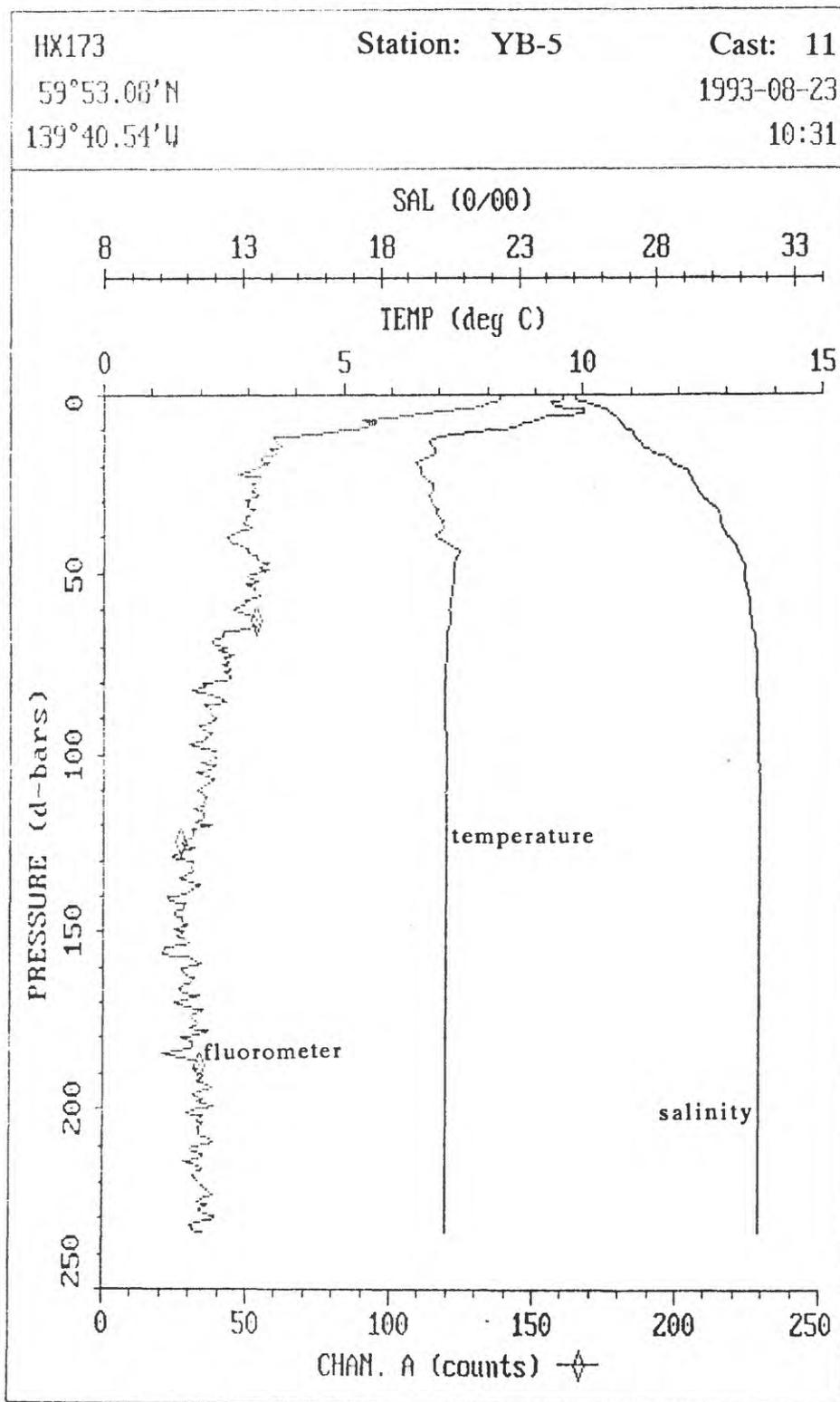


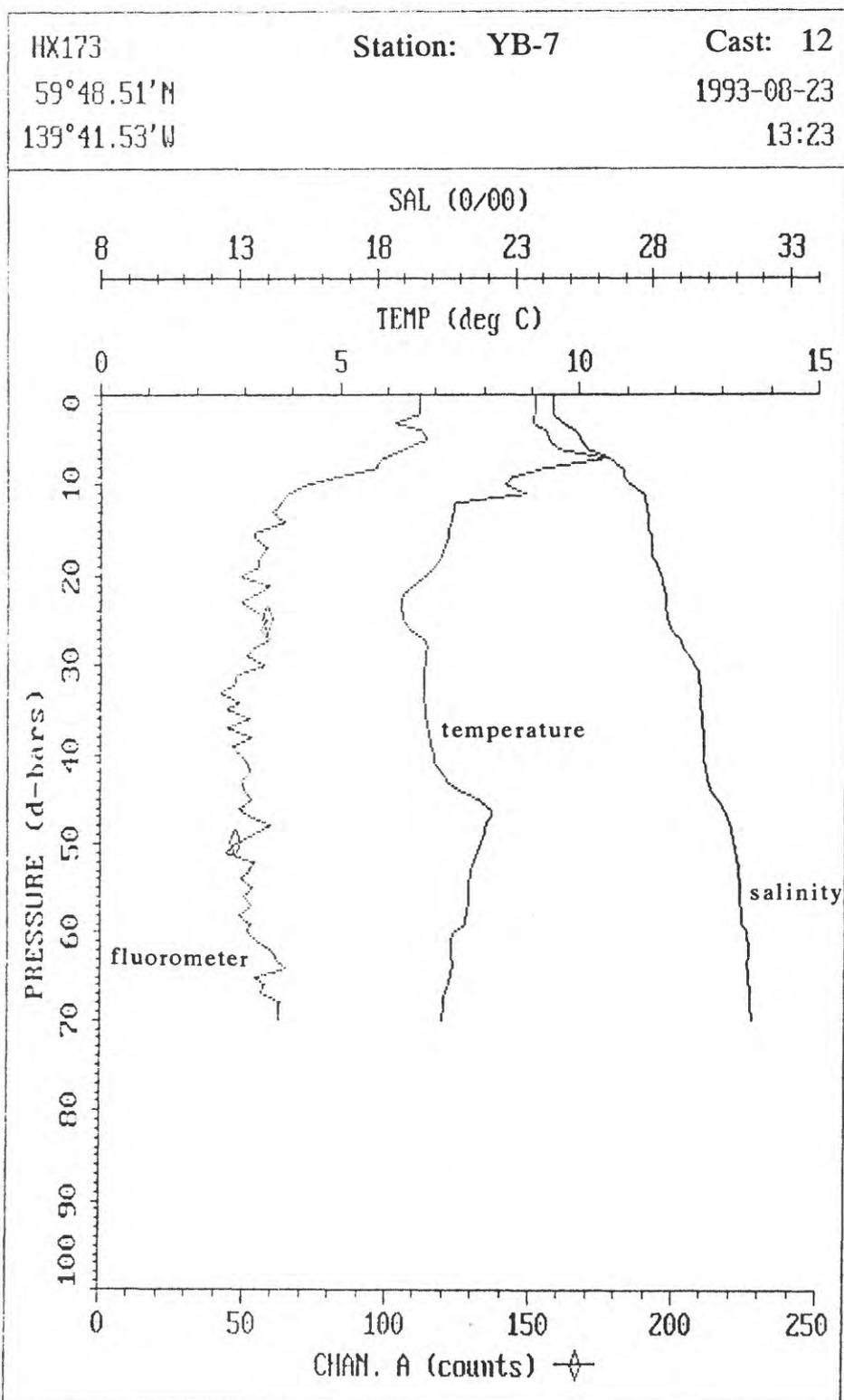


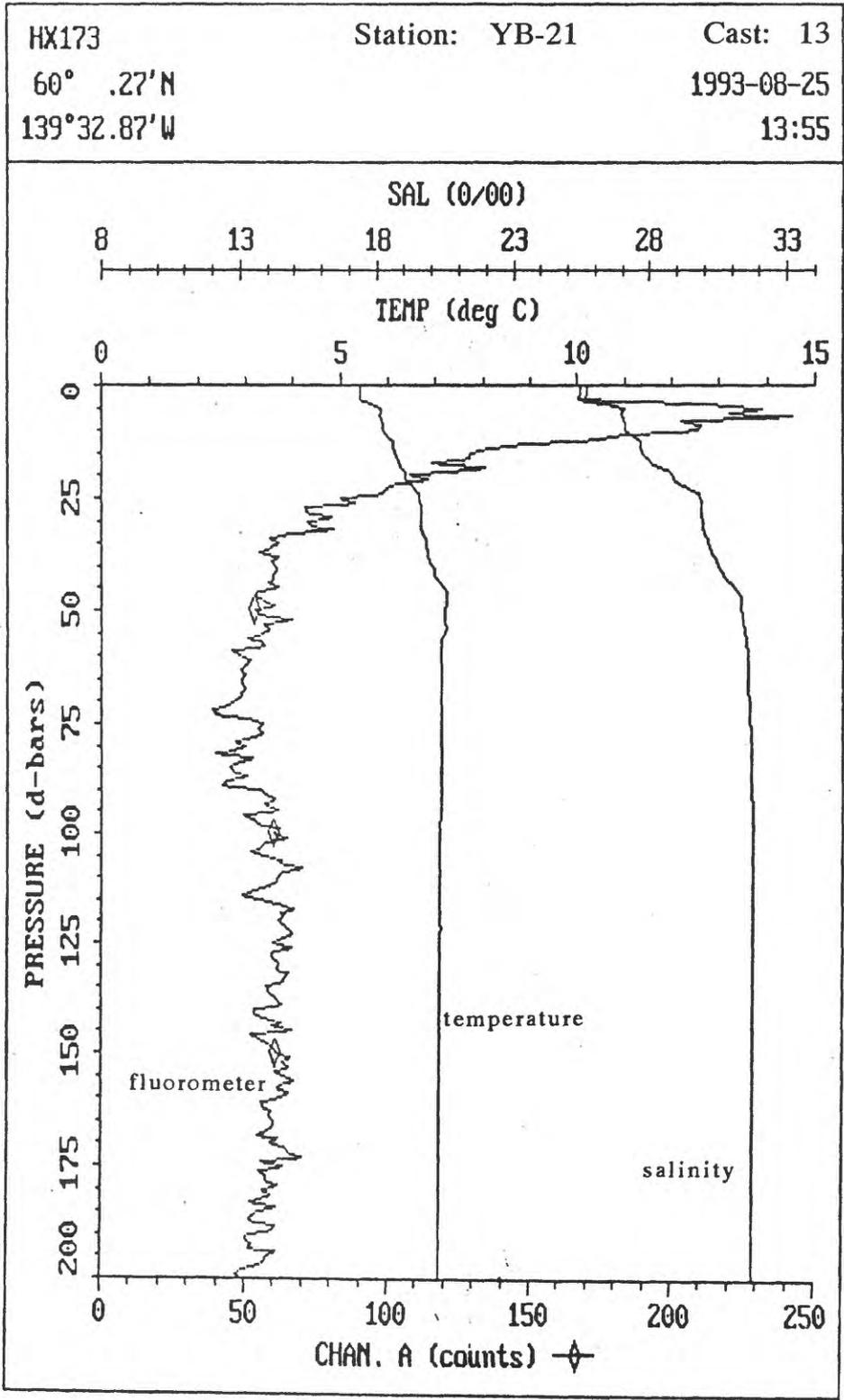


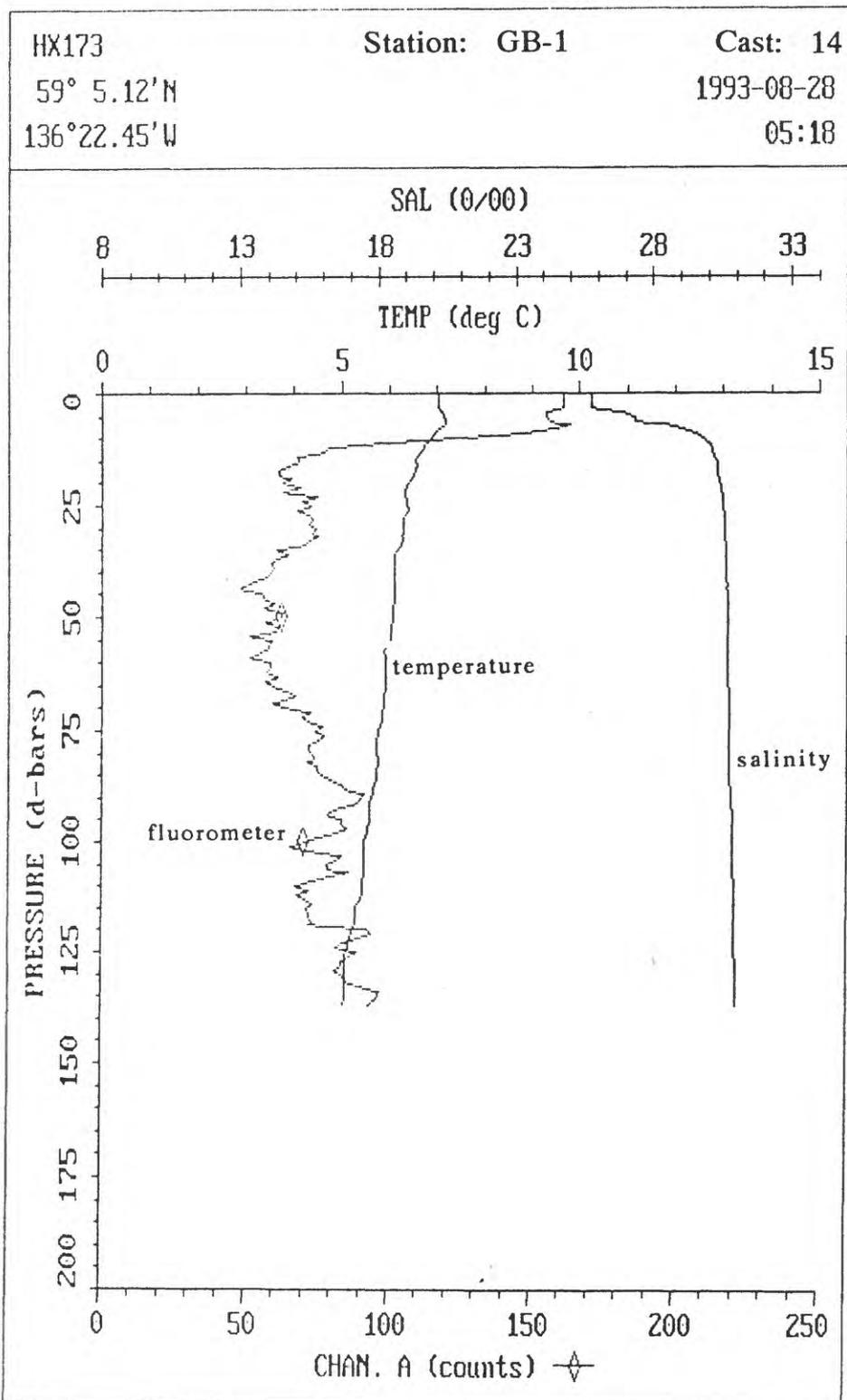


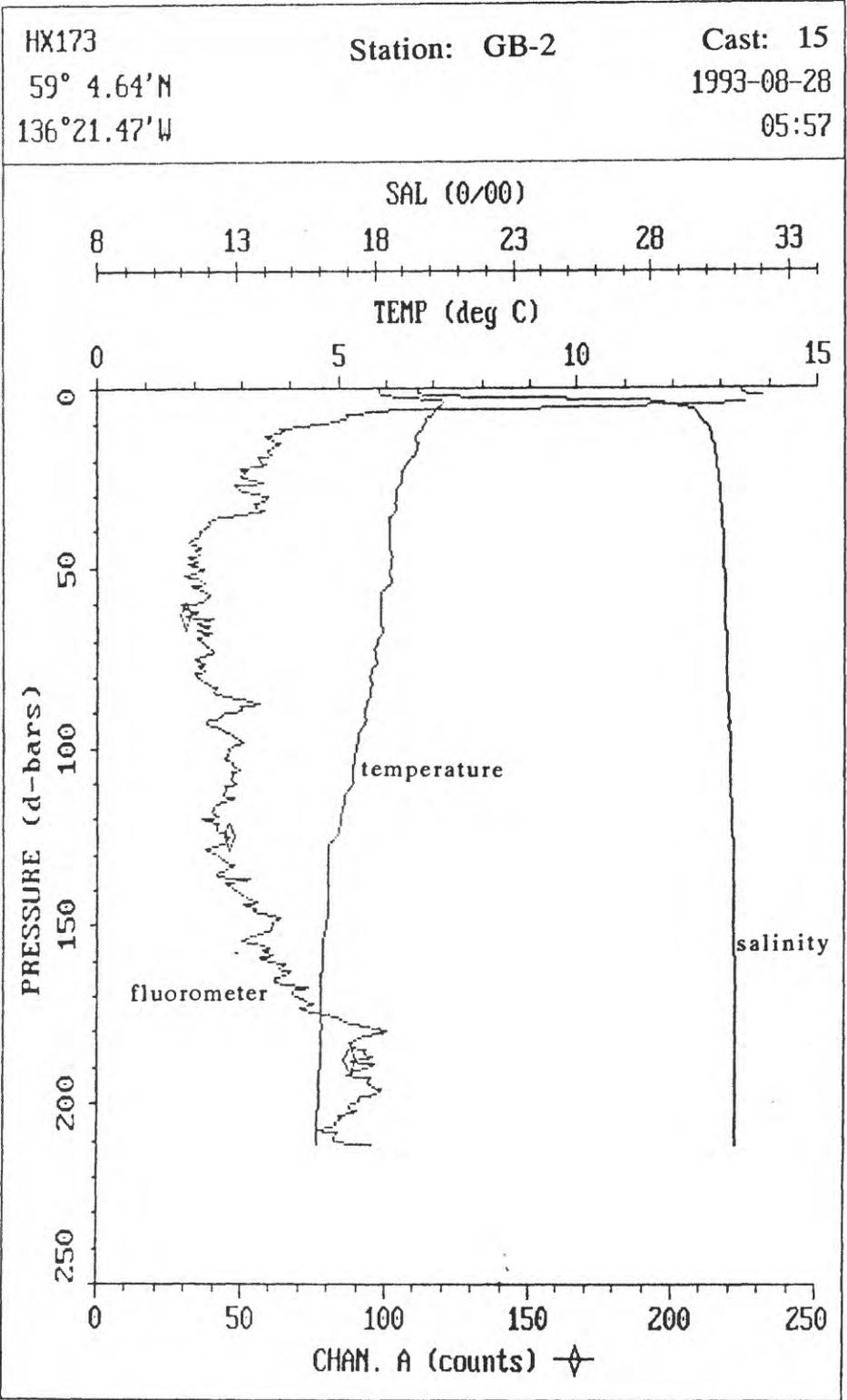




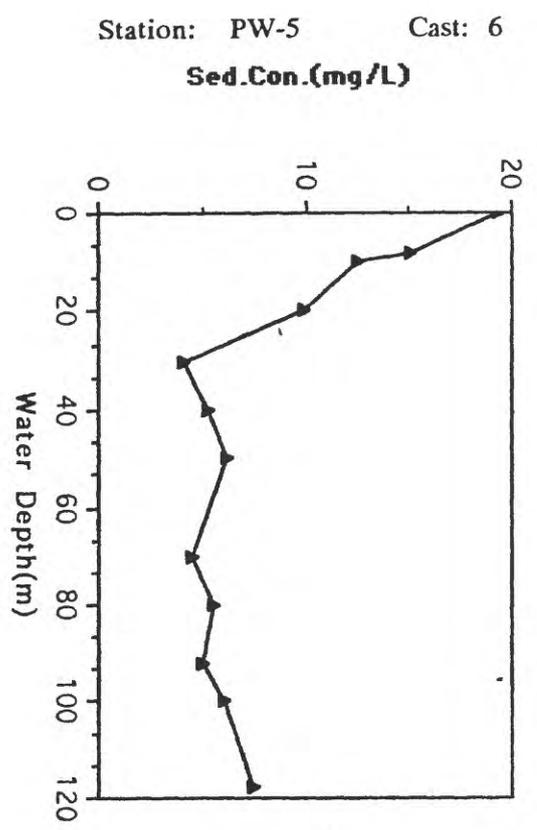
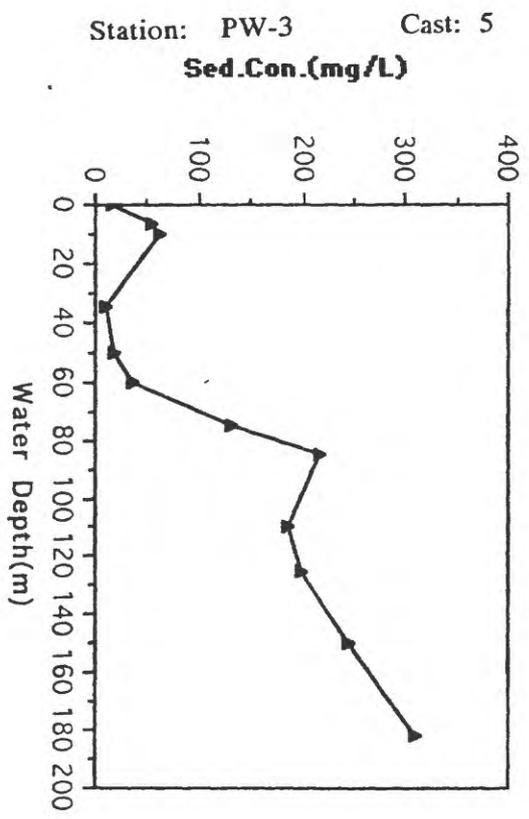
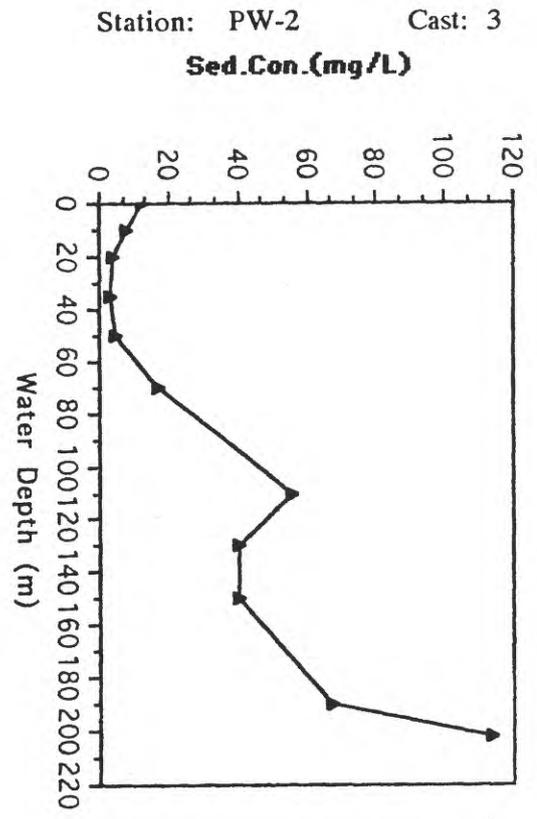
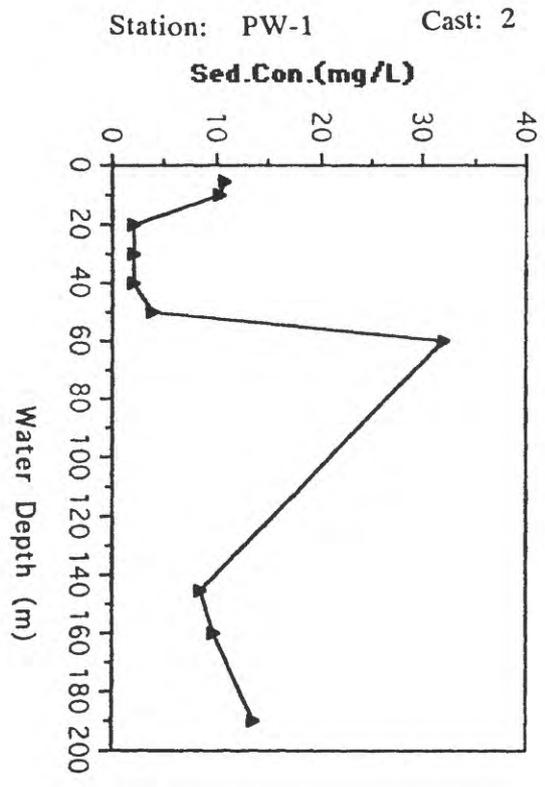






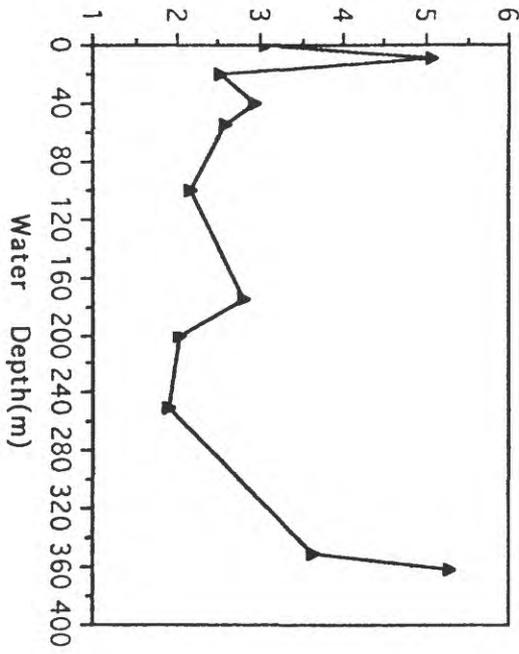


**APPENDIX 4: PLOTS OF SUSPENDED
SEDIMENT CONCENTRATIONS**



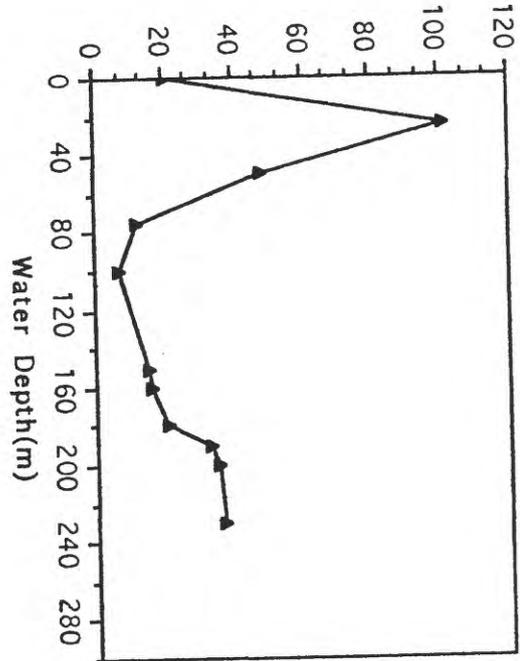
Station: PW-6 Cast: 7

Sed.Con.(mg/L)



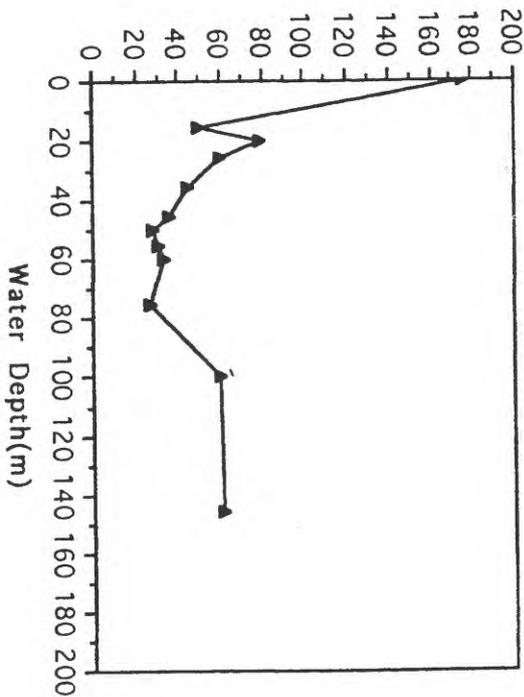
Station: YB-1 Cast: 8

Sed.Con.(mg/L)



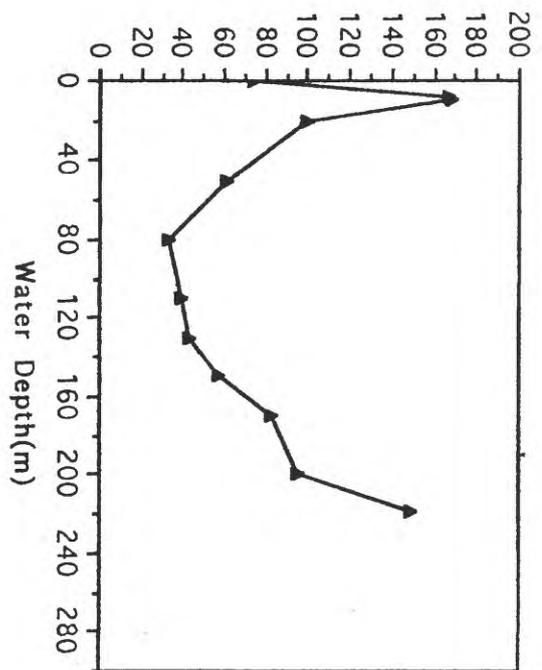
Station: YB-2 Cast: 9

Sed.Con.(mg/L)

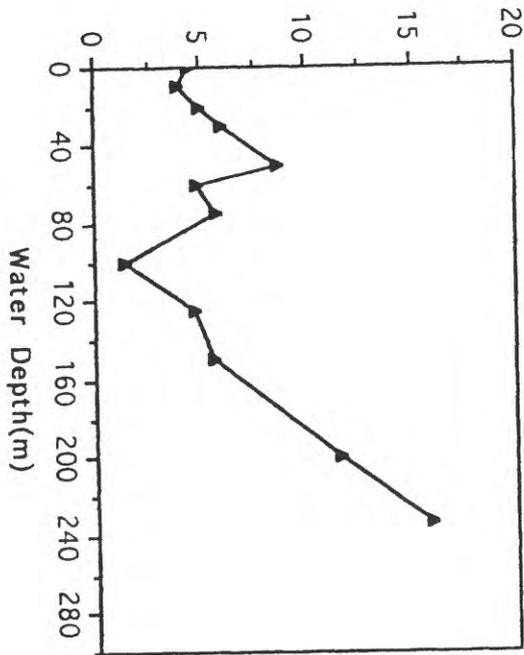


Station: YB-4 Cast: 10

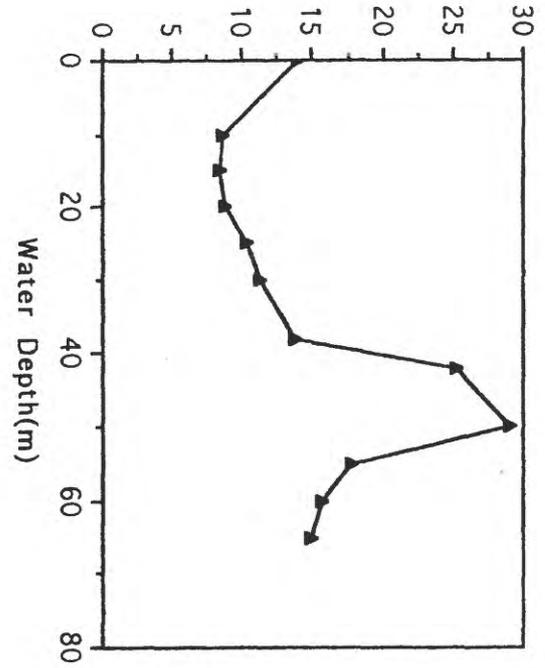
Sed.Con.(mg/L)



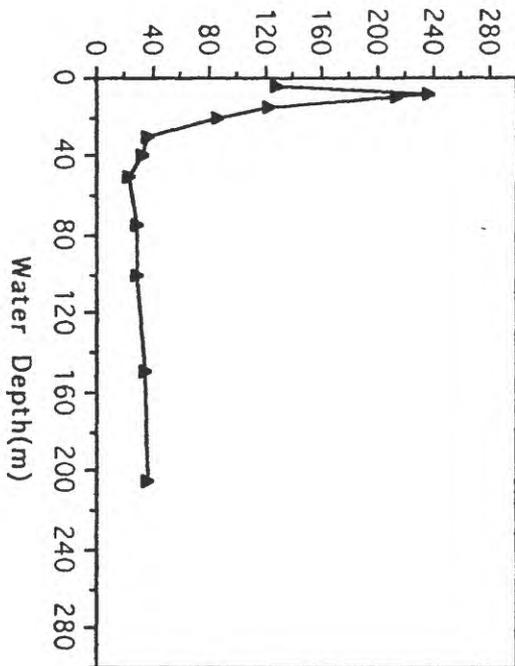
Station: YB-5 Cast: 11
Sed.Con.(mg/L)



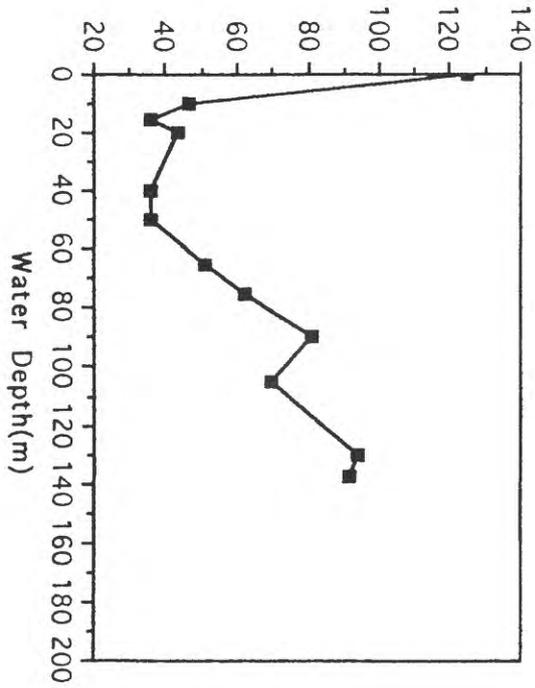
Station: YB-7 Cast: 12
Sed.Con.(mg/L)



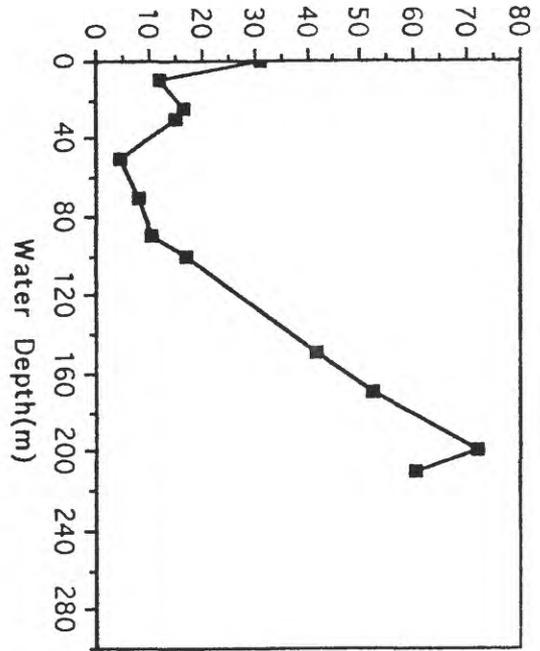
Station: YB-21 Cast: 13
Sed.Con.(mg/L)



Station: GB-1 Cast: 14
Sed.Con.(mg/L)



Station: GB-2 Cast: 15
Sed.Con.(mg/L)



Station: GB-34 Cast: 16
Sed.Con.(mg/L)

