

Acknowledgments

In 1991, the USGS began collection of bathymetric data for the reaches in the GIS data base to extend topographic information to the channel bottom to define the channel morphology for model development. Reach 5, which is about 19 km long and extends from river mile 60 to river mile 72² (see index map below), was selected as the reach in which to begin model development because the Little Colorado River, which joins the Colorado River at about river mile 61, is the largest source of sand to the Colorado River in the park. The morphology of sand deposits in the reach downstream from the confluence is important to native fish, especially the endangered humpback chub, because areas associated with sandbars—backwaters, return-flow channels, channel margins—are habitats for

chub hatched in the warmer waters of the Little Colorado River (Angradi and others, 1992, Valdez and others, 1992). Also, the reach includes 32 monumented cross sections established in 1992 and 1993 to monitor changes in channel sand storage. Cross sections in the A, B, and C series (sheets 2 and 3) were established and measured for the first time in June and July 1992, and those in the D, E, and F series (sheets 4–7) in February 1993. Methods of data collection and analysis for the cross-section measurements are described by Graf and others (1995).

These maps present the combined topographic and bathymetric data used to develop grids of equally spaced points representing channel morphology required for the development of multidimensional models of flow, sediment transport, and bed evolution. The topography and bathymetry are presented in sheets 2-7, and data extend from just below the Little Colorado River confluence to just above Tanner Rapids—10.3 km of reach 5 from about river mile 62 to about mile 68.5. Bathymetry in the remaining parts of reach 5 had not been mapped at the time of preparation of this report. The BOR produced 8 map sheets of the topography of reach 5. Data presented in this report cover sheets of BOR map sheet 3 (sheet 2) and 8 (sheet 7) and all of map sheets 4-7 (sheets 3-6).

Data used to produce the maps have been used to generate two-dimensional arrays of equally spaced elevations (grids) used to provide channel geometry for multidimensional models of flow, sediment transport, and bed evolution (Wiele, S.M., Graf, J.B., and Smith, J.D., research hydrologists, USGS, written commun., 1995). The maps also serve as base maps for precise location of measurements and sample locations with respect to important geomorphic features.

² Locations of features along the Colorado River in the study reach typically are given in miles downstream from Lees Ferry, Arizona. River miles are used in this report for consistency with local convention.

Depth was measured with a sonic depth sounder, and the digital position and depth data were sent directly or by radio modems to a datalogger or laptop computer. Depth also was recorded on paper charts. The two positioning systems were used interchangeably. The accuracy of the two systems was determined by a specified accuracy of ± 0.5 percent ± 0.025 m of indicated depth. The unit sounds the depth about nine times per second and was used with a transducer that had a beam width of 8° . The unit used in 1993 had a manufacturer's specified accuracy of ± 0.5 percent ± 0.025 m of indicated depth. The unit used in 1994 had a transducer beam width was 9° . Both units were calibrated in the field before the surveys by suspending a lead weight or plate at known depths that ranged from about 2 to 20 m below the transducer. Adjustments were made to the speed-of-sound value and the depth of the water with the recorded depth to the depth of weight at the deepest calibration depth.

Verification of Bathymetric Data

Bathymetric surveys of 10.3 km of reach 5 were made in June and July 1992 and April and May 1993 (see index map below). Bathymetric data were tied to the Arizona State Plane Coordinate System using a network of control points. The control-point locations were determined by USGS surveys and verified by other GCES surveys. During the bathymetric surveys, the position of a target mounted to the top of the depth-sounder transducer on the boat was measured and combined with depth data measured at the same time.

The area covered by the topographic data is largely bedrock, talus, and gravel in cliffs, gravel bars, debris fans, and talus slopes. These features were not altered significantly by river flows that occurred from the time of the aerial photograph until after the second bathymetric survey. The area covered by the bathymetric surveys is subject to greater and more frequent changes than is the area covered by the topographic data.

A flood of moderate peak discharge but long duration—3 weeks—occurred on the Little Colorado River in January 1993, between the two bathymetric surveys. The flood was caused by a snowmelt event in the headwaters of the study area and caused aggradation of the bed at many of the monumented cross sections in the A, B, and C series (sheets 2 and 3), which were the only sections measured before and after the flood (Jain and others, 1994). Graf and coauthors (1994) estimated that the flood caused aggradation of the bed at 15 of the 19 cross sections. The bed aggradation in January 1993 was initially deposited in the vertical and lateral channel extensions typically found in pools just downstream from the constrictions formed by debris fans (see cross sections 1000 and 1001). The bed aggradation was not measured in the S.M., Graf, J.B., and Smith, J.D. research hydrologists, USGS, writing commun.,

The digital data collected in the field were edited to identify and remove data with incorrect boat position and (or) incorrect depths. Points with incorrect positions were found through the use of computer programs that flagged points more than a selected distance from points collected previously and by inspection of plotted values. Large sets of points with incorrect positions and individual points with incorrect positions and depths were deleted from the data set. When small sets of points with incorrect positions but correct depths were found, the data information was retained and new positions were calculated by linear interpolation between the two positions on either end of the gap along the boat track. Over distances of a few meters, the interpolated positions are within the accuracy of the bathymetric survey.

Incorrect depths are found by comparison of the digital record with the paper chart record collected at the same time. The digital recorder stores only the first signal returned, but the paper-chart record typically contains a record of the channel bottom as well as any material within the water column that reflects the signal (air bubbles, algae, and (or) fish). Where the channel bottom was recorded on the paper chart but the digital record contained incorrect depths, the paper chart record was digitized and the incorrect depths were replaced with the digitized depths. The data were deleted in a few cases for which the bottom was not recorded on either the paper chart or the digital recorder.

After verification of position and depth data from the bathymetric survey, the positions of the data points were converted from arbitrary x,y coordinates to a system in which they were collected into the Arizona State Plane Coordinate system. System used in the 1992 data base is Wertz and others, 1992. The 1992 bathymetric survey depth data were converted to water-surface-elevation data collected at the time of the survey. Each pool reach was surveyed separately and a single water-surface elevation for the pool was used to convert the data. The water-surface elevation in the pool was determined from depth measured by a temporary gauge that consisted of a pressure transducer and an recorder. Depth data were related to a datum of known elevation. The water-surface elevation was determined by adding the depth to the datum elevation. The bathymetric data. Because the adjustment was made on a pool-by-pool basis, the water-surface elevation change during each survey was not significant. Although during the plan had been to make the adjustments for the surveys done in 1992 in the same manner as was done for the 1993 surveys, uncertainties in the contours from the survey done at the time of the 1992 bathymetric survey made the water-surface elevation adjustment for the 1992 survey data difficult. The water-surface elevation was adjusted by adding the data from each pool survey until the best match in elevation with the topographic data was obtained. The best match was determined by visual comparison of elevation at points in the bathymetric data set with the topographic contours. This method proved satisfactory because the topographic data were obtained at a lower discharge than the bathymetric surveys, and the bathymetric surveys were obtained in many areas especially in gently sloping areas such as debris fans.

Data from the two bathymetric surveys were combined with topographic data for contouring and plotting. Because of real differences in elevation that were a consequence of channel changes over time and apparent differences caused by position uncertainty, points in areas of overlap between the topographic data and the bathymetric data sometimes differed significantly in elevation. In those cases points from the bathymetric set in the area of overlap were deleted because those points had a greater position uncertainty.

To generate contours, a network of triangles (TIN) was created from the combined topographic set using the Delaunay method of triangulation (McCullagh and Ross, 1980) implemented by the ARC/INFO software. Because the density of data points from the bathymetric surveys along traverse lines created a point distribution that produced unrealistic contours, the unrealistic contours—within a given radius of the previous point—were smoothed into sharp peaks and narrow small-scale features were suppressed by the smoothing process. The reason the TIN contained many extremely long narrow triangles (point spacing of 4 km was determined best for these data because the

triangles. A point spacing of 4 ft was determined best for these data because the spacing produced a TIN with the best geometry—closest to a network of equilateral triangles—and contours that were best supported by the data. An outer boundary close to but outside of the outer points was used to define the external boundary of the network. Once the TIN was developed, an interpolation algorithm using a bivariate fifth-degree polynomial in x and y (Akima, 1978) was used to compute contour-line positions from the network. Contours were evaluated for accuracy by comparison to contours developed photogrammetrically for the topographic data set. Resulting contours closely match the original topographic

contours (see inset on sheet 3). In some places, including areas of very steep or shallow slope and areas of sparse data, unrealistic or incorrect contours were generated. These contours were edited to better represent the real topography.

A contour interval of 1 m was used wherever feasible. On sheets 2 and 3, a contour interval of 5 m was used for contours above 830 m because the extremely steep slopes in those areas made the 1-meter contours unreadable. In some areas, the contours at the outer edge of the map end abruptly and appear to be cut off. In these areas, the topographic data ended abruptly, and no information was

available to generate contours beyond these points. In most cases, these abruptly terminated contours abut near-vertical slopes, and these areas are indicated by cliff lines on the map.

Contour maps (sheets 2-7) were developed from combined bathymetric and topographic data for the Colorado River in a 10.3-kilometer reach downstream from its confluence with the Little Colorado River.

The positions of monumented cross sections on the maps, and the cross-section geometry for a single survey at each of these sections is shown to illustrate the channel geometry. Cross sections shown may not agree exactly with the geometry shown by the contours because the bathymetric surveys and cross-section measurements have different accuracies and were made at different times.

Bathymetry of the area upstream from a point 0.7 km upstream from Carbon Creek (sheet 4) was mapped before the Little Colorado River flood of January 1993, and the area downstream from that point was mapped 3–4 months after the flood. The index map below and on each sheet shows the area and dates of the two bathymetric surveys, and the boundary between the surveys is shown on sheet 4.

The surveyed area is indicated by the location of data points from the bathymetric surveys used to generate the contours. Areas in the channel not covered by surveys were either very shallow or are areas of rapids or riffles where water velocity and turbulence was too high for the depth sounder to return a bottom signal. To help the reader visualize the channel, the extent of the water surface at a discharge of about 142 m³/s in 1989 also is shown. The water surface data were obtained by the BOR from aerial photographs taken at the discharge and are shown on sheets 2-7 without modification. River miles also are shown on the maps because they are a commonly used system of documenting locations along the river.

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Multiply metric unit	<u>B₂</u>	To obtain inch-pound unit
meter (m)	3.28	foot
cubic meter per second (m ³ /s)	35.3	cubic foot per second
kilometer (km)	0.621	mile
teragram (Tg)	1.103 × 10 ⁶	short ton

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