

# Geodesy and Contemporary Strain in the Yucca Mountain Region, Nevada

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## CONVERSION FACTORS, VERTICAL DATUM, AND ABBREVIATIONS

	Multiply	By	To obtain
centimeter (cm)		0.3937	inch
meter (m)		3.281	foot
kilometer (km)		0.6214	mile
millimeter (mm)		0.03937	inch
square kilometer (km <sup>2</sup> )		0.3861	square mile

## ADDITIONAL ABBREVIATIONS

g	gravitational acceleration
mm/yr	millimeter per year
Ma	Millions of years
VLBI	Very Long Baseline Interferometry

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## Abstract

Geodetic surveys provide important information for estimating recent ground movement in support of seismotectonic investigations of the potential nuclear-waste storage site at Yucca Mountain, Nevada. Resurveys of established level lines document up to 22 millimeters of local subsidence related to the 1992 Little Skull Mountain earthquake, which is consistent with seismic data that show normal-slip rupture and with data from a regional trilateration network. Comparison of more recent surveys with a level line first established in 1907 suggests 3 to 13 centimeters of subsidence in the Crater Flat-Yucca Mountain structural depression that coincides with the Bare Mountain fault; small uplifts also were recorded near normal faults at Yucca Mountain. No significant deformation was recorded by a trilateration network over a 10-year period, except for coseismic deformation associated with the Little Skull Mountain earthquake, but meaningful results are limited by the short temporal period of that data set and the small rate of movement. Very long baseline interferometry that is capable of measuring direction and rates of deformation is likewise limited by a short history of observation, but rates of deformation between 8 and 13 millimeters per year across the Basin and Range province are indicated by the available data.

## INTRODUCTION

Since the late 1970s, Yucca Mountain, in southwestern Nevada, has been investigated by the U.S. Department of Energy in cooperation with agencies including the U.S. Geological Survey, as a potential national geologic repository for the storage of high-level radioactive wastes. As part of a broad interdisciplinary program which is designed to evaluate comprehensively the suitability of the Yucca Mountain site (U.S. Department of Energy, 1988), a series of specific studies that bear on the structural geology, tectonic evolution, and seismicity of the Yucca Mountain region (fig. 1), including the one discussed herein, was undertaken in the mid-1980s to analyze ground motion and fault-displacement hazards that may affect repository design and performance. Site-characterization efforts generally have focused on a specific site area (fig. 2) in which mapping and related studies were conducted in greater detail than in surrounding areas. A fan-shaped "controlled area" of some 90 square kilometers (fig. 2) results from regulatory requirements (Title 10 of the Code of Federal Regulations, 10CFR60 and 10CFR960) and surrounds a "potential repository site" that represents the upward projection of the boundary of the proposed underground storage facilities.

## Background and Purpose

This report is one of a series that describes results and presents interpretations of available information on the tectonics and seismicity of Yucca Mountain, including those geologic elements (structural, stratigraphic, volcanic, and geomorphic) that contribute to the makeup and tectonic evolution of

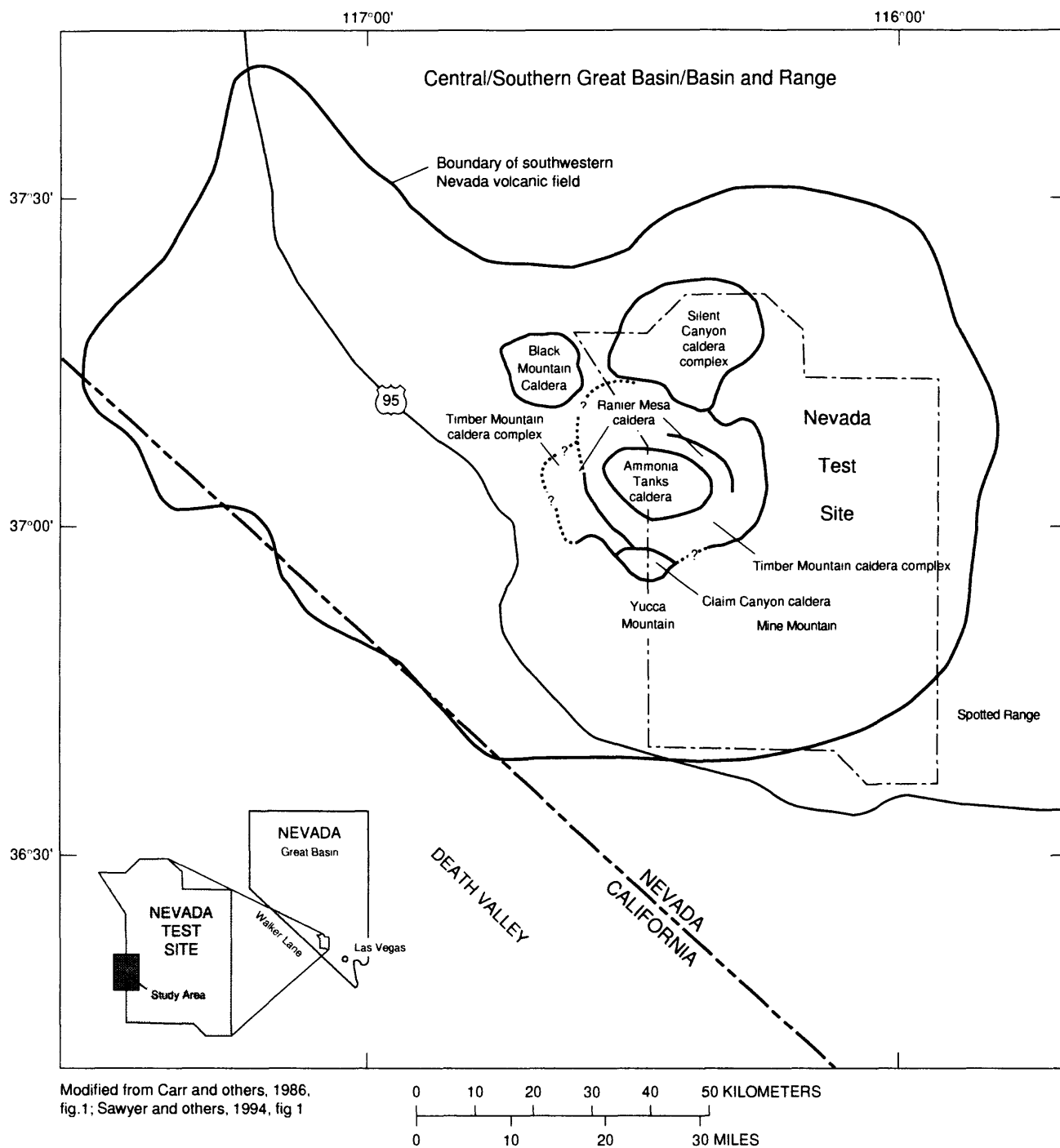


Figure 1. Locations of the southwestern Nevada volcanic field and selected volcanic and physiographic features of the Yucca Mountain region. The study area is also shown as "site area" on figure 2. Modified from Carr and others, 1986, and Sawyer and others, 1994.

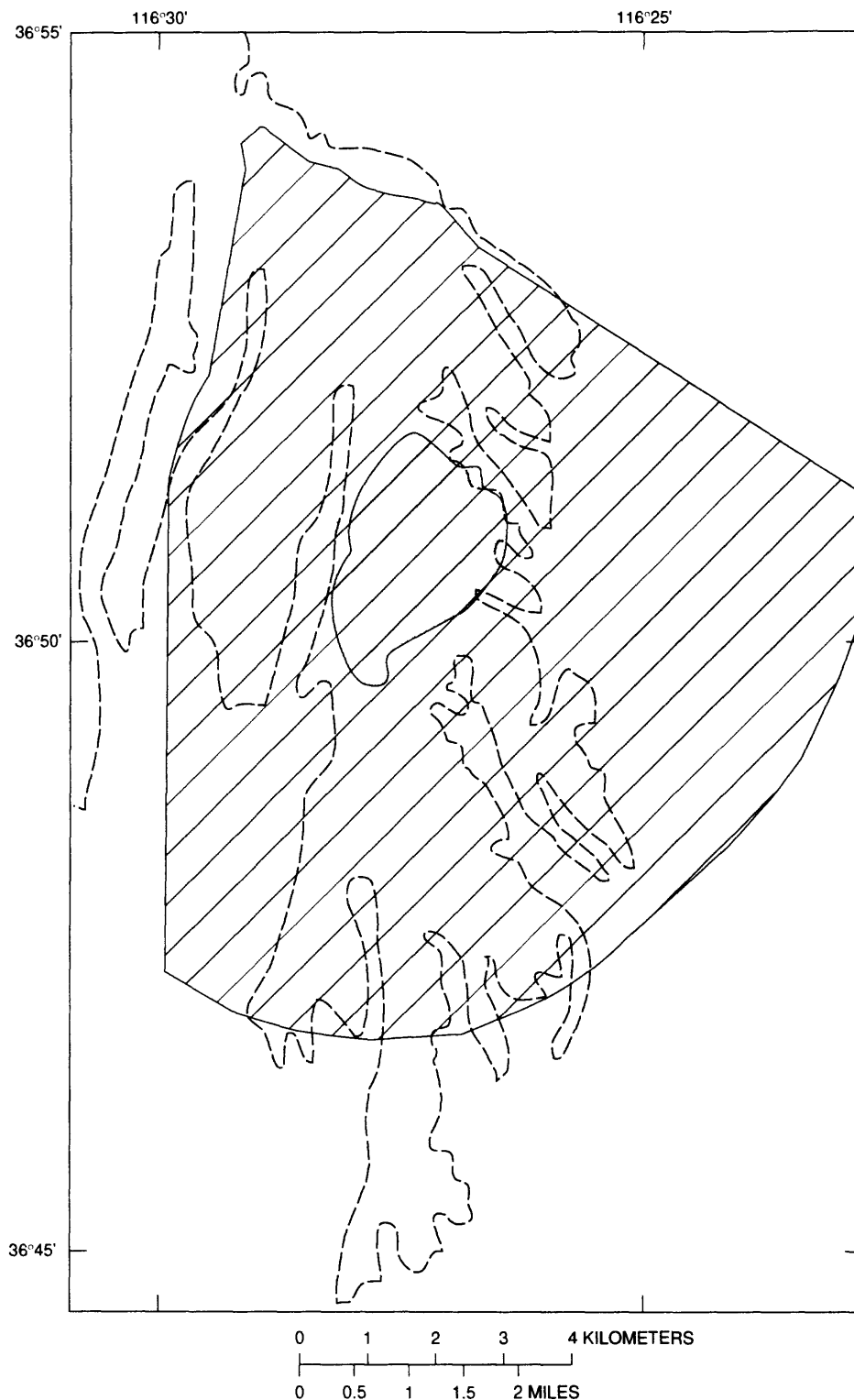


Figure 2. Index map of Yucca Mountain (dashed outline) and vicinity, Nevada, showing boundaries of areas commonly referred to in site-characterization studies and reports. Outside boundaries of the map define the "site area" (same area as shown in figure 1) encompassing most of the significant geologic features being mapped and studied in detail at or in close proximity to Yucca Mountain. The extent of the potential repository (upward projection of the boundaries of the underground facility) is shown by the shaded area, and the extent of the controlled area is marked by diagonal lines. Modified from Rautman and others, 1987.

Yucca Mountain, neighboring ranges and bounding basins, the proximal fault zones that distinguish the structural pattern of Yucca Mountain and surrounding areas (typical of the southern Great Basin), the foci of seismic events, and adjacent volcanic fields. Of most interest to the site-characterization program are features that generally occur within 100 km of Yucca Mountain (fig. 1).

Subjects covered in these reports were discussed in detail in the Yucca Mountain Site Characterization Plan (SCP; U.S. Department of Energy, 1988). Techniques and methods of study supporting specifically defined objectives are described in widely reviewed study plans for individual investigations. The scope of these structural, seismic, and tectonic investigations covered (1) identification and characterization of faults, both at Yucca Mountain and in surrounding areas, with emphasis on faults with demonstrable Quaternary activity; (2) distribution and geochronology of Quaternary deposits, data that record paleoseismicity and lead to calculations of amounts, rates, and recurrence intervals of Quaternary fault displacements; (3) seismicity in the southern Great Basin and earthquake magnitude and recurrence rates at Yucca Mountain; (4) location and characterization of relevant earthquake sources; and (5) development of tectonic models.

Regulations embodied in the Nuclear Waste Policy Act of 1982 and the Code of Federal Regulations (10CFR60 and 10 CFR960) require evaluation of natural phenomena that might compromise the design, safety, and performance of a geologic repository, including potentially adverse conditions such as landslides, subsidence, volcanic eruptions, changes in climatic conditions and water-table levels, structural deformation (especially during the Quaternary), and seismic activity and frequency. Specific information needs include identification of relevant earthquake sources, estimation of peak acceleration, amounts of ground motion and other ground-motion characteristics, preparation of data sets for seismic-source evaluation, consideration of historical seismicity, characterization of Quaternary faults within 100 km of Yucca Mountain, mapping and analysis of fault zones, determination of the extent of detachment faulting in the region, and geodetic leveling to determine contemporary deformation. The combined information provides the basis for the development of tectonic models that may contribute to a better and more

comprehensive understanding of the processes that could impact repository design and performance.

Rates of tectonic processes during recent, historical, and pre-historical Quaternary times—the three separate time periods of particular significance to Yucca Mountain site characterization—are determined from contemporary observations, recorded history, and geochronological data on Quaternary deposits and fault movements, respectively. This report summarizes results of various geodetic surveys that have been conducted in the Yucca Mountain area and the surrounding region to provide quantitative data on crustal movements and strain accumulations during the most recent of these time periods.

## Geologic Setting

Yucca Mountain, which is located near the western boundary of the Nevada Test Site in southwestern Nevada about 140 km northwest of Las Vegas, lies in the north-central part of the Basin and Range physiographic province and consists of a series of rugged north-trending fault-block ridges composed of volcanic rocks that generally have an eastward tilt of 5° to 10° (Scott and Bonk, 1984). The potential site area is situated on the southern flank of the Miocene southwestern Nevada volcanic field (Byers and others, 1976; fig. 1). A thick section of Tertiary volcanic rocks (table 1) overlies Paleozoic sedimentary strata in the Yucca Mountain region (Byers and others, 1976), with units of the 12.8 to 12.7 Ma Paintbrush Group (Sawyer and others, 1994) forming most exposures (Christiansen and Lipman, 1965; Scott and Bonk, 1984). Two voluminous densely welded ash-flow tuffs of the Paintbrush Group underlie Yucca Mountain (the Tiva Canyon and Topopah Spring Tuffs), separated by a much thinner interval of mostly nonwelded pyroclastic rocks (Buesch and others, 1996); the tuffaceous rocks are greater than 3,000 m thick in some places (Snyder and Carr, 1984). High-angle faults locally exhibit displacements of several hundred meters (Carr, 1984; Scott and Bonk, 1984). The geometry of these normal faults and of northwest-striking strike-slip faults controls most of the topography and drainage at Yucca Mountain. The right-lateral strike-slip faults have been linked to the Walker Lane (Carr, 1984), a large structural zone (fig. 1) that parallels the southwestern border of Nevada. Extensional faulting is present between the

Paleozoic and Tertiary rocks (Carr and Monsen, 1988) at the northern end of Bare Mountain (fig. 1), an uplifted fault block consisting of complexly deformed sedimentary and metasedimentary rocks of Paleozoic and Precambrian age. Bare Mountain bounds the western side of Crater Flat, a 200-km<sup>2</sup> depositional basin adjacent to Yucca Mountain (fig. 3). This alluvial flat is the most strongly extended and most deeply subsided part of the Crater Flat structural basin, a graben-like domain that also includes Yucca Mountain (Fridrich, in press).

**Table 1.** Major stratigraphic units in the Late Cenozoic volcanic rock sequence of the Yucca Mountain area, Nevada. Ages are given in millions of years (Ma). Modified from Sawyer and others (1994) and Buesch and others (1996).

Unit	Age (Ma) <sup>1</sup>
Timber Mountain Group	
Ammonia Tanks Tuff	11.45
Rainier Mesa Tuff	11.6
Post-Tiva Canyon pre-Rainier Mesa rhyolites <sup>2</sup>	12.5
Paintbrush Group	
Tiva Canyon Tuff	12.7
Yucca Mountain Tuff	--
Pah Canyon Tuff	--
Topopah Spring Tuff	12.8
Calico Hills Formation	12.9
Crater Flat Group	
Prow Pass Tuff	--
Bullfrog Tuff	13.25
Tram Tuff	--
Lithic Ridge Tuff	14.0

<sup>1</sup>Ages have not been determined for all units.  
<sup>2</sup>This informal grouping represents multiple rhyolites of small areal extent erupted from multiple sources at the margins of calderas such as the Timber Mountain caldera complex north of Yucca Mountain. Correlations are incompletely understood for all such identified rhyolites and calderas.

Many studies (summarized in Volume 1, Part A, of the SCP for Yucca Mountain; U.S. Department of Energy, 1988, section 1.3) have shown that the tectonic setting of Yucca Mountain is complex, with the area situated within the Basin and Range physiographic province near the intersection of the northwest-trending Walker Lane and the northeast-trending Spotted Range-Mine Mountain structural zone (for discussion, see Carr, 1984). The

Walker Lane is a major structural lineament and appears to be a zone of transition between an area to the north and east characterized by dip-slip (normal) faulting, and an area to the south and west characterized by both dip-slip and right-lateral strike-slip faulting. Well-exposed low-angle detachment faults have been documented in regions that surround Yucca Mountain (Hamilton, 1988; Maldonado, 1990; Scott, 1990; Simonds and others, 1996; and Hoisch and others, 1997). Those detachment faults and the closely spaced high-angle faults at Yucca Mountain led some investigators to propose detachment(s) beneath Crater Flat and Yucca Mountain (Hamilton, 1988; Spengler and Fox, 1989; and Scott, 1990), but recent geophysical results are inconclusive on this issue (Brocher and others, 1996; Majer and others, 1996).

### GEODETIC LEVELING

Determination of contemporary motion on geologic structures requires establishment of a network of level lines from which to determine movement. An existing first-order level line extends from a point on U.S. Highway 95 (benchmark location S16, fig. 3) north-northeast across Crater Flat, east over Yucca Mountain, southeast across Jackass Flats and between Skull and Little Skull Mountains, across Rock Valley and ends back on Highway 95 (benchmark location I17, fig. 3). This 92-km-long line, first established and surveyed during 1956–59, was designed to cross the Rock Valley fault and the major structural blocks and block-bounding faults of Yucca Mountain. It contains 133 benchmarks, one per kilometer except for a 20-km interval across Yucca Mountain, where spacing shortens to about 0.5 km. Since 1983, the line has been resurveyed every 1 or 2 years (1983, 1983–84, 1985–86, 1988, 1990–91, and 1992–93); recordings are contained in unpublished geodetic leveling reports by T. K. Bray (U.S. Geological Survey, written commun., 1993), G. C. Perasso (U. S. Geological Survey, written commun., 1993), and E. I. Balazs (National Geodetic Survey, written commun., 1993). Survey observations made after 1984 meet first-order Class 1 standards established by the Federal Geodetic Control Committee (1984), and the 1990–91 and 1992–93 series were double-run surveys.

A preliminary analysis of the level-line survey data was made primarily to examine possible effects



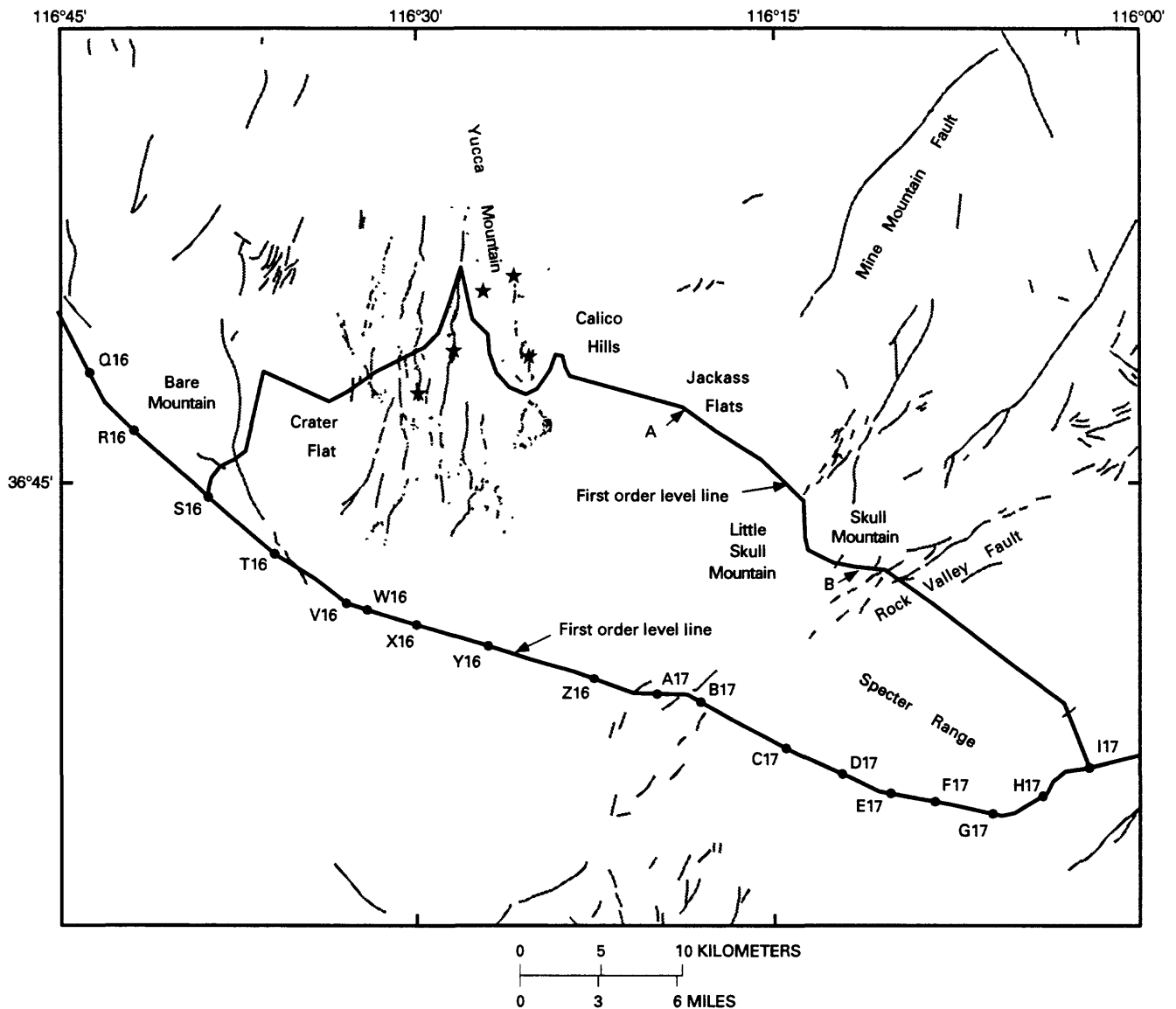


Figure 3. Index map showing location of first-order level lines in the vicinity of Yucca Mountain, Nevada. Points A and B are in reference to figure 4, and benchmark locations (for example, W16) are in reference to figure 5. Faults shown are modified from Piety (1996) and Simonds and others (1995). Stars show locations of quadrilaterals across major faults at Yucca Mountain.

of the June 29, 1992, Little Skull Mountain earthquake (M 5.8) and also to determine whether there had been any noticeable change in benchmark elevations across Yucca Mountain during 1983–93. The analysis consisted of a comparison of the pre- and post-earthquake benchmark elevations by (1) subtracting the post-earthquake elevations from the pre-earthquake elevations, and (2) plotting the elevation differences with respect to a reference line that represented values recorded during the 1985–86 survey (K.S. Koepsell, National Geodetic Survey, written commun., 1996). The results are shown in figure 4. Based principally on data from the 1983, 1985–86, 1988, and 1992–93 surveys, the most noticeable observation to be made is that the pre-earthquake benchmark elevations in the vicinity of Little Skull Mountain were higher than the post-earthquake elevations (fig. 4). The negative elevation changes recorded in the 1992–93 survey occur in an area that lies between 53 and 70 km east of the west end of the level line (points A and B in fig. 3), with the maximum change of 22 mm observed at a distance of about 62 km, which is 2 km northwest of Little Skull Mountain. The downdropped area coincides closely with an area that lies between the southwestward projection of the Mine Mountain fault in Jackass Flats and the Rock Valley fault in Rock Valley (fig. 3).

The observed decrease in benchmark elevations in the vicinity of Little Skull Mountain is in close agreement with the interpretation of K.D. Smith and others (University of Nevada/Reno, written commun., 1995) who described the main shock of the Little Skull Mountain earthquake as involving primarily down-to-the-southeast dip-slip movement on a steeply dipping, northeast-striking fault plane. Meremonte and others (1995) reported that (1) a projection of the mainshock fault plane to the surface, delineated by its focal mechanism and distribution of aftershocks, defines a strike subparallel to a north-northeast structural trend within Jackass Flats and collinear with the Mine Mountain fault system; and (2) the mainshock focal mechanism and the aftershock distribution indicate a normal-slip rupture. Savage and others (1994) also found that normal slip provided a close fit to the resulting deformation observed in a 50-km-aperture trilateration network centered on Yucca Mountain (see discussion below). It is interesting to note that the Eureka Valley earthquake (M 6.1) that occurred May 17, 1993, on the border between California and Nevada resulted in a depression similar in width and

depth to that created by the Little Skull Mountain event (Peltzer and Rosen, 1995).

The correlation of elevation change with topography indicates that there may have been elevation-dependent errors along some sections of the level line, thus calling into question whether or not actual deformation is being recorded. This relation is most apparent across Yucca Mountain (from 25 to 35 km), particularly in the elevations recorded during the 1983, 1983–84, and 1988 surveys. Also noticeable, but less apparent, is the correlation of elevation change with topography toward the southeastern end of the survey line (75 to about 95 km). The pronounced deviations of the recordings made during the 1983–84 and 1990–91 surveys from the more consistent values obtained in the other surveys suggest that further processing and analysis are required to determine the cause(s) of the deviations and the overall validity of the survey results.

Level-line surveys also have been conducted along Highway 95 between Beatty and Las Vegas, beginning with a third-order leveling survey by the U.S. Geological Survey in 1907, followed by first-order surveys by the National Geodetic Survey in 1915 and again during the 1980–84 period. The accumulated geodetic data have been compiled by Gilmore (1992), with special attention given to changes in elevation at bench-mark locations spaced 1.5 to 6 km apart on that part of Highway 95 that is between the western and eastern ends of the level line across Yucca Mountain (fig. 3). By using the 1907 survey as a baseline, the 1915 and 1980–84 data show negative height changes ranging from 3 to 13 cm for an area south of Bare Mountain and Crater Flat (between bench marks Q16 and V16, figs. 3 and 5), which is referred to as the Crater Flat/Yucca Mountain structural depression by Gilmore (1992). The Bare Mountain fault zone coincides with the largest of these negative height changes. Eastward from bench mark W16 (fig. 3), positive height changes of 5 to 8 cm were recorded along a portion of the level line that is located directly south of mapped segments of several of the north-striking normal faults at and near Yucca Mountain (fig. 5). Height changes increase still farther eastward across the southwestward projection of the Rock Valley fault zone, leveling off to a relative uplift of 15 to 20 cm over the broad region between that point and the Las Vegas area. In general, the 1980–84 survey shows negative height changes with respect to the 1915 survey.

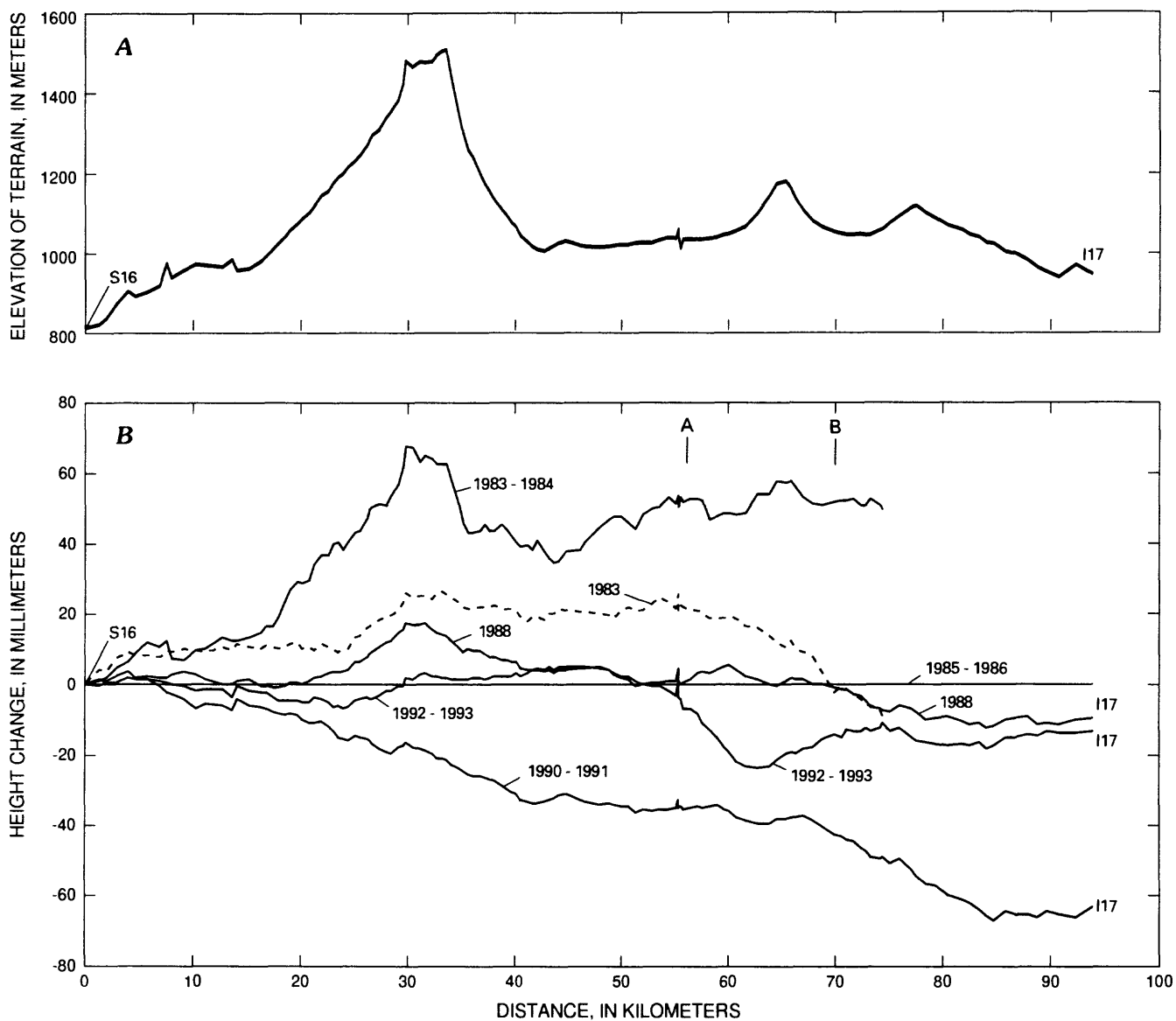


Figure 4. A., Profile plot showing data from surveys of level line across Yucca Mountain, Nevada, during the period 1983-1993 (after K.S. Koepsell, National Geodetic Survey, written commun., 1996); B., Data from the 1985-1986 survey are used as the baseline to show changes in elevation. Location of level line points A and B are shown on figure 3.

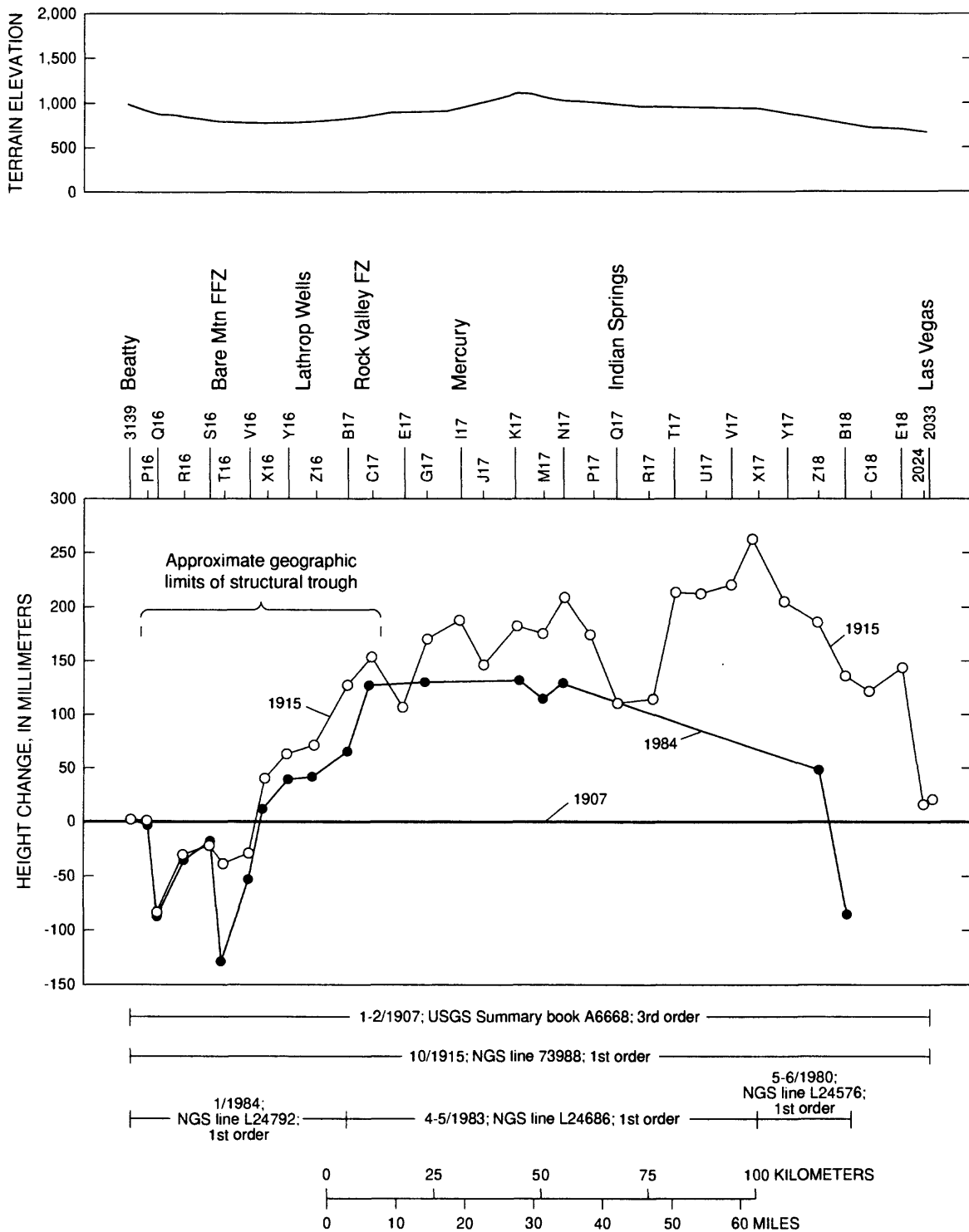


Figure 5. Profiles showing terrain elevations and height changes with respect to the 1907 baseline between Beatty and Las Vegas, Nevada (from Gilmore, 1992). Locations of bench marks Q16 eastward to I17 are shown on figure 3.

Quadrilateral networks, consisting of four corner stations a few hundred to several hundred meters apart and arranged in near-rectangular patterns, were established across several of the major normal faults at Yucca Mountain in the early 1980s (fig. 3). The distances between the stations and their elevations were measured in 1983, 1983–84, 1984, 1985–86, and 1988 with surface-based surveying instruments. During this 5-year period, no changes in either the horizontal or vertical positions of the stations in any of the quadrilaterals were detected that could be interpreted as providing clear evidence of recent fault movement (J.M. Vukovich, U.S. Geological Survey, oral commun., 1996). Beginning in the early 1990s, the quadrilaterals were resurveyed using Global Positioning System (GPS) methods; GPS surveys over a period of several more years are needed before meaningful results can be obtained because the rate of movement, if there is any, is so small that it cannot be identified over a short period of time within the limits of error of the methods involved.

## TRILATERATION - GLOBAL POSITIONING SYSTEM

Deformation within a 50-km-aperture trilateration network in the Yucca Mountain area has been periodically measured by the U.S. Geological Survey during 1983–93 (Savage and others, 1994; see fig. 6). The network was surveyed in 1983 and 1984 by measuring distances between stations with a geodolite, whereas in 1993 the distances were surveyed using GPS, and some were measured with both geodolite and GPS to provide data for correcting differences between the two methods. The resulting survey data were interpreted by Savage and others (1994) to indicate that no significant deformation was detected across the trilateration network for this 10-year period, except for coseismic deformation associated with the June 29, 1992, Little Skull Mountain earthquake in the southeastern corner of the network (fig. 6). Using a fault plane determined from seismic data (dip 54°SE, strike N35°E), Savage and others (1994) determined that  $0.580 \pm 0.075$  m of normal slip on a 5-km<sup>2</sup> rupture surface at a depth of about 8 km provided a close fit to the observed deformation in that part of the network.

Horizontal and vertical changes in the positions of about 20 stations are being recorded by GPS methods along a 140-km-long profile that stretches

from Beatty, Nevada, southwest to Coso Junction, California. This line spans several major structural features, including the Furnace Creek fault zone, Death Valley, the Panamint Mountains, the Panamint Valley fault zone, and prominent faults near Coso Junction. Results of these surveys are not yet available.

A regional network of GPS stations was established jointly by the Nuclear Regulatory Commission, California Institute of Technology, and the Smithsonian Institute in 1991 (Southwest Research Institute, 1995). The network consists of 15 survey sites in 3 subnetworks: Yucca Mountain — 5 sites to measure strain across the Bare Mountain fault zone, Crater Flat, Yucca Mountain proper, and Jackass Flats; Death Valley — 4 sites to determine the rate of strain accumulation across the Northern Death Valley-Furnace Creek fault zone, plus 2 additional sites in the central part of Death Valley to provide control for determining strains across the Death Valley-Furnace Creek and Hunter Mountain-Panamint fault zones at positions south of the Death Valley and Hunter Mountain subnets; and Hunter Mountain — 4 sites for measuring strain across the Hunter Mountain fault (see fig. 7 for fault locations). Preliminary results of GPS surveys conducted in 1991, 1993, and 1994 indicate two important trends in relation to contemporary strain patterns in the Yucca Mountain region, including (1) vertical and horizontal motions around the potential radioactive waste repository site may be present and would be detectable with further monitoring; and (2) significant horizontal (as much as 5 mm/yr) and vertical (as much as 20 mm/yr) strains are indicated along major fault zones within 50 to 100 km of Yucca Mountain, but accumulated motions are barely measurable by the 1991–94 surveys (Southwest Research Institute, 1995). There is uncertainty regarding strain rates because of the short period of time represented by these surveys. Continued monitoring of the network, however, should bring improvement in these uncertainties.

## VERY LONG BASELINE INTERFEROMETRY

Very Long Baseline Interferometry (VLBI) is a relatively new geodetic technique that allows measurements of the directions and rates of change of position of stations that are typically separated by

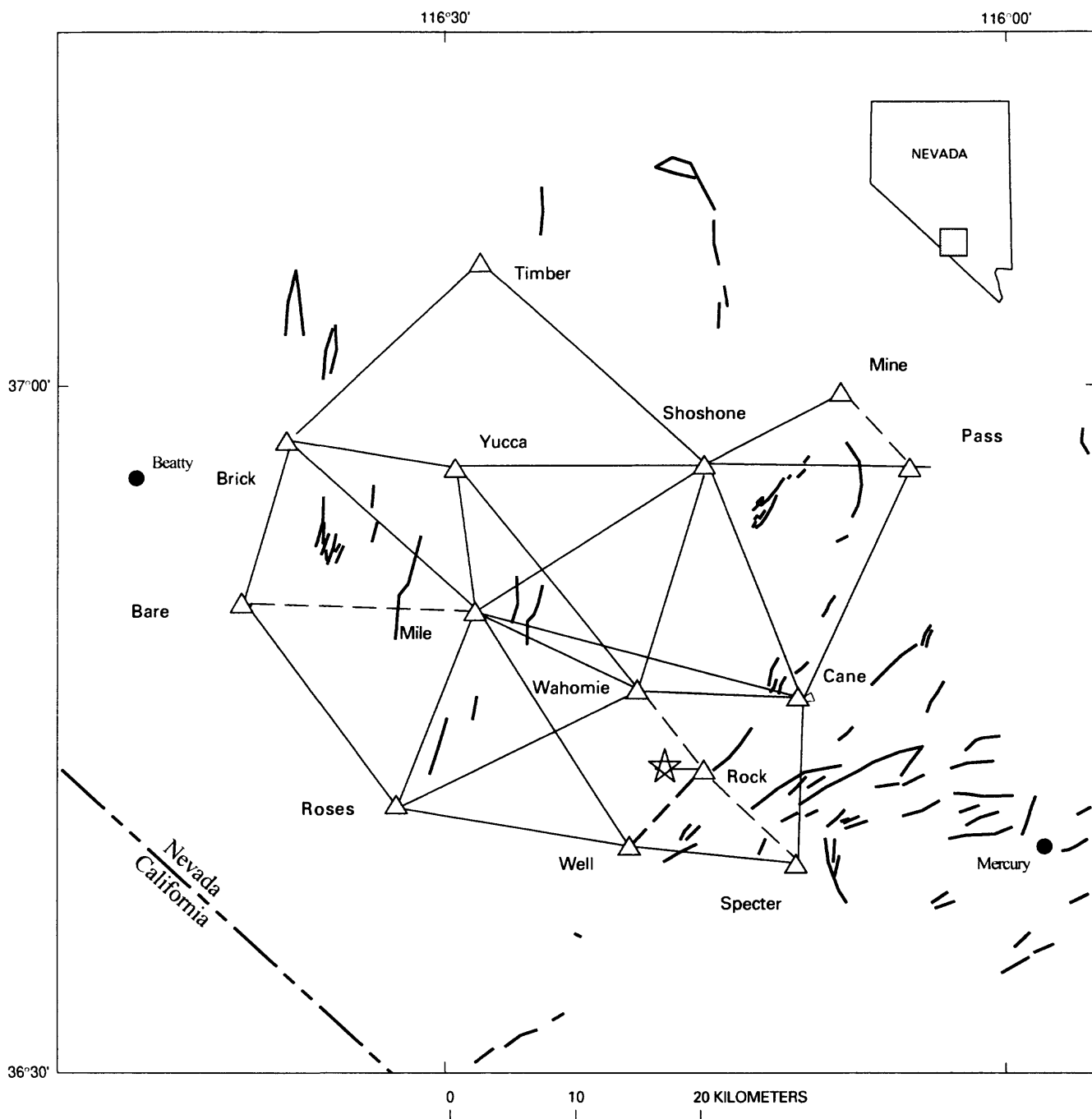
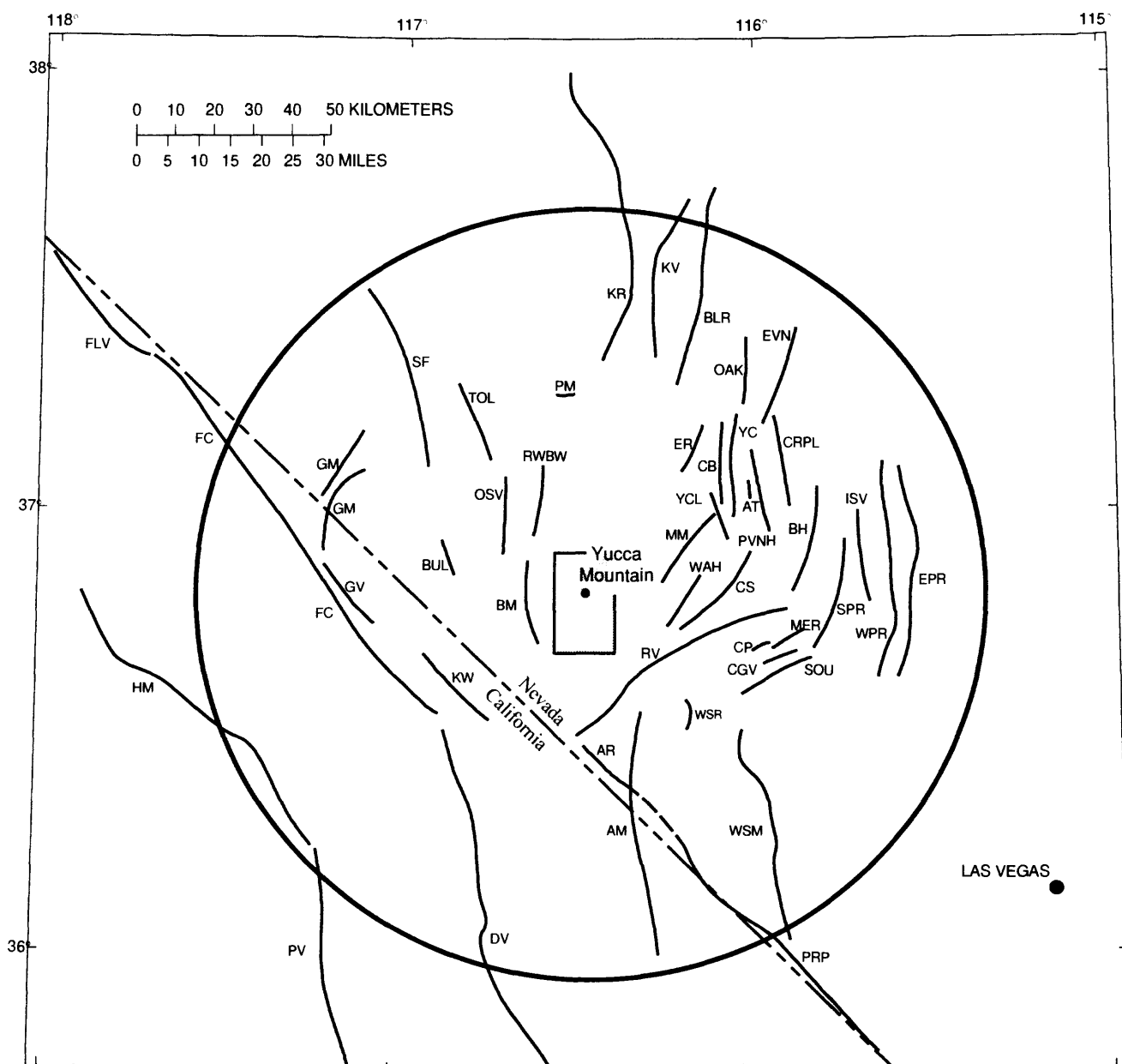


Figure 6. Trilateration network near Yucca Mountain, Nevada. Lines show the measured distances; solid lines were measured in all three surveys, but dashed lines were included in only one of the 1983 and 1984 surveys. The star is the epicenter of the 1992 Little Skull Mountain earthquake. The potential Yucca Mountain waste disposal site is at station Mile. Sinuous lines represent faults. (Modified from Savage and others, 1994).



EXPLANATION  
Abbreviations of faults

AM	Ash Meadows	CS	Cane Spring	KR	Kawich Range	RV	Rock Valley
AR	Amargosa River	DV	Death Valley	KV	Kawich Valley	RWBW	Rocket Wash-Beatty Wash
AT	Area Three	EPR	East Pintwater Range	KW	Keane Wonder	SF	Sarcobatus Flat
BH	Buried Hills	ER	Eleana Range	MER	Mercury Ridge	SOU	South Ridge
BLR	Belted Range	EVN	Emigrant Valley North	MM	Mine Mountain	SPR	Spotted Range
BM	Bare Mountain	FC	Furnace Creek	OAK	Oak Spring Butte	TOL	Tolicha Peak
BUL	Bullfrog Hills	FLV	Fish Lake Valley	OSV	Oasis Valley	WAH	Wahmonie
CB	Carpetbag	GM	Grapevine Mountains	PM	Pahute Mesa	WPR	West Pintwater Range
CGV	Crossgrain Valley	GV	Grapevine	PRP	Pahrump	WSM	West Springs Mountain
CP	Checkpoint Pass	HM	Hunter Mountain	PV	Panamint Valley	WSR	West Specter Range
CRPL	Cockeyed Ridge-Papoose Lake	ISV	Indian Springs Valley	PVNH	Plutonium Valley-N Halfpint Range	YC	Yucca
						YCL	Yucca Lake

Figure 7. Index map showing location of relevant and potentially relevant faults capable of generating average median and 84<sup>th</sup>-percentile peak accelerations that equal or exceed 0.1 g at the potential waste repository at Yucca Mountain, Nevada. The large circle is 100 km in radius from the site. To preserve the sense of the number of faults in the region, this index figure shows many faults not mentioned in the text.

200 km or more. Thus, VLBI data provide a picture of deformation over scales of several hundred kilometers, even distances as long as across the entire Basin and Range Province, but observations are temporally limited to slightly more than the last decade. VLBI data from across the western United States have been interpreted to indicate that, overall, the Basin and Range is deforming at rates between 8 and 13 mm/yr, in a northwesterly direction (N30°W to 45°W), relative to stations on and near the Colorado Plateau (Sauber, 1989; DeMets and others, 1987, 1990; Ward, 1990; Argus and Gordon, 1991; Dixon and others, 1995). Dixon and others (1995) suggested that this motion can be resolved into 5 mm/yr east-west extension (presumably on the Wasatch fault) east of Ely, Nevada, and about 10 mm/yr northward motion of the Sierra Nevada with respect to Ely. Geologic, smaller scale geodetic, and seismicity studies indicate that three zones along the western edge of the southern Great Basin — the Owens Valley, Panamint Valley-Hunter Mountain, and Death Valley-Furnace Creek-Fish Lake Valley fault zones — accommodate most of the deformation that extends northward from the Mojave Desert along what has been referred to as the Eastern California Shear Zone (Dokka and Travis, 1990; Savage and others, 1990; Richard, 1992; Dixon and others, 1995; Reheis and Dixon, 1996). Based on combined geologic and geodetic studies in the Death Valley region, Reheis and Dixon (1996) suggested that several of the major northwest-striking fault zones are connected by short north-striking normal faults that transfer slip among the major faults via right steps. At present, the total slip rate of  $10.7 \pm 1.6$  mm/yr (Dixon and others, 1995) across the shear zone near latitude 36°15' is distributed among the Owens Valley fault zone, the Hunter Mountain-Panamint Valley faults, and the Furnace Creek fault zone, but northward the slip is increasingly transferred eastward to the Fish Lake Valley fault zone (Reheis and Dixon, 1996). The eastern limit of the Eastern California Shear Zone is recognized to be the Death Valley, Furnace Creek, and Fish Lake Valley fault zones, which are located approximately 50 km west-southwest of Yucca Mountain at their closest proximity (see fig. 7 for fault locations).

## CONCLUSIONS

The geodetic surveys show no changes in elevation that can confidently be attributed to fault movements at and near Yucca Mountain, except in response to the June 29, 1992, Little Skull Mountain earthquake. In that case, both the resurveys of the level line in the vicinity of Little Skull Mountain and the measurements obtained from a trilateration network that spans the Yucca Mountain area are in agreement with the amount and type of motion deduced from seismic records. Regional geodetic surveys show that strain and large slip rates exist within 50 to 100 km west of Yucca Mountain, but the motions are only barely resolvable by GPS methods and the current temporally limited data sets.

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