

EVALUATION OF SPECIFIC-STORAGE VALUES

Calibrated models of aquifer-system compaction developed by Helm (1975, 1976, 1977, 1978) and by Williamson and others (1989) indicate that a narrow range of both aquitard inelastic skeletal specific storage values and aquitard elastic skeletal specific storage values have been used to successfully simulate aquifer-system compaction despite differences in depth, thickness, lithology, and stratigraphy among the San Joaquin Valley sites that were modeled (table 4). The aquitard inelastic skeletal specific storage values derived by Helm (1978) for seven sites in the San Joaquin Valley ranged from about 1.4×10^{-4} to $6.7 \times 10^{-4} \text{ ft}^{-1}$, and have a mean of about $3.2 \times 10^{-4} \text{ ft}^{-1}$ and a standard deviation of about $1.8 \times 10^{-4} \text{ ft}^{-1}$. The aquitard elastic skeletal specific storage values that Helm (1978) derived ranged from about 2.0×10^{-6} to $7.5 \times 10^{-6} \text{ ft}^{-1}$, and have a mean of about $4.5 \times 10^{-6} \text{ ft}^{-1}$ and a standard deviation of about $2.1 \times 10^{-6} \text{ ft}^{-1}$. Aquitard elastic and inelastic skeletal specific storage values used by Williamson and others (1989) fall within the ranges of values derived by Helm.

The narrow range of aquitard inelastic and elastic skeletal specific storage values derived by Helm (1978) at these seven sites also has been found elsewhere in the United States, including areas of California other than the San Joaquin Valley. For seven sites in the Santa Clara Valley, California, model-derived values of aquitard inelastic skeletal specific storage ranged from 1.4×10^{-4} to $4.0 \times 10^{-4} \text{ ft}^{-1}$ (Helm, 1978) and have a mean of $2.8 \times 10^{-4} \text{ ft}^{-1}$ and a standard deviation of $8.9 \times 10^{-5} \text{ ft}^{-1}$. The model-derived values of aquitard elastic skeletal specific storage for the seven Santa Clara Valley sites ranged from 2.2×10^{-6} to $1.6 \times 10^{-5} \text{ ft}^{-1}$ (Helm, 1978) and have a mean of $6.7 \times 10^{-6} \text{ ft}^{-1}$ and a standard deviation of $5.2 \times 10^{-6} \text{ ft}^{-1}$.

At an extensometer site in Antelope Valley in the Mojave Desert of southern California, the model-derived value of aquitard inelastic skeletal specific storage for aquitards that were actively compacting inelastically was $3.5 \times 10^{-4} \text{ ft}^{-1}$ (Sneed and Galloway, 2000), which is strikingly close to the mean value of $3.2 \times 10^{-4} \text{ ft}^{-1}$ for the 7 sites in San Joaquin Valley. At this same Antelope Valley site, the model-derived value of aquifer-system elastic skeletal specific storage was $1.7 \times 10^{-6} \text{ ft}^{-1}$ (Sneed and Galloway, 2000). This value is slightly smaller than those reported for the San Joaquin and Santa Clara Valleys, perhaps because the elastic skeletal specific storage value for the Antelope

Valley site represents an average value for the aquifer system, rather than an elastic skeletal specific storage value that explicitly represents the aquitard component of the aquifer system.

Hanson (1989) reported model-derived estimates for selected extensometer sites in the Tucson basin and in the Avra Valley using the Helm model approach. For the six Arizona sites, the model-derived values of aquitard inelastic skeletal specific storage ranged from 7.3×10^{-6} to $2.7 \times 10^{-4} \text{ ft}^{-1}$ (Hanson, 1989) and have a mean of $9.4 \times 10^{-5} \text{ ft}^{-1}$ and a standard deviation of $1.1 \times 10^{-4} \text{ ft}^{-1}$. Smaller values in this range may indicate that the aquifer systems were still in transition to predominantly inelastic compaction when the study was done; hence long-term inelastic skeletal specific-storage values may initially increase because of increased compaction through time (Hanson, 1989). The model-derived values of aquitard elastic skeletal specific storage for the six Arizona sites ranged from 1.0×10^{-6} to $2.0 \times 10^{-5} \text{ ft}^{-1}$ (Hanson, 1989) and have a mean of $7.1 \times 10^{-6} \text{ ft}^{-1}$ and a standard deviation of $6.6 \times 10^{-6} \text{ ft}^{-1}$. Epstein (1987) reported on model-derived estimates for an extensometer site near Eloy, Ariz. The model-derived value of aquitard inelastic skeletal specific storage for the upper five layers was $1.5 \times 10^{-4} \text{ ft}^{-1}$; layers 6 and 7 used 1.8×10^{-4} and $2.7 \times 10^{-4} \text{ ft}^{-1}$, respectively. The model-derived value of the aquitard elastic skeletal specific storage for layers 1 through 5 and 7 was $3.0 \times 10^{-6} \text{ ft}^{-1}$; layer 6 used $2.4 \times 10^{-6} \text{ ft}^{-1}$ (Epstein, 1987).

Inelastic skeletal specific storage measured from laboratory consolidation tests on samples tend to result in smaller values compared with those derived from calibrated models of aquifer-system compaction. At site 14S/13E-11D, the model-derived estimate of inelastic skeletal specific storage was $4.3 \times 10^{-4} \text{ ft}^{-1}$ (Helm, 1978). For the same location, laboratory consolidation tests on eight samples yielded inelastic skeletal specific storage values ranging from 2.0×10^{-5} to $1.6 \times 10^{-4} \text{ ft}^{-1}$ (Johnson and others, 1968) (table 3). The larger value was measured in a sample from the Corcoran Clay, which would be expected to be highly compressible, yet this value is smaller than Helm's (1978) derived value by a factor of about 2.7. At site 16S/15E-34N, the model-derived estimate of inelastic skeletal specific storage was $2.4 \times 10^{-4} \text{ ft}^{-1}$ (Helm, 1978). For the same location, laboratory consolidation tests on 16 samples yielded inelastic skeletal specific storage values ranging from 1.4×10^{-5} to $1.7 \times 10^{-4} \text{ ft}^{-1}$

(Johnson and others, 1968) (table 3); the largest value in this range is smaller than Helm's (1978) by a factor of about 1.4. In this example, the largest values were not measured in samples from the Corcoran Clay, but in sediments collected above the Corcoran Clay (Johnson and others, 1968) (table 3). At site 23S/25E-16N, the model-derived estimate of inelastic skeletal specific storage was $2.3 \times 10^{-4} \text{ ft}^{-1}$ (Helm, 1978). For the same location, laboratory consolidation tests of four samples yielded inelastic skeletal specific storage values ranging from 4.0×10^{-5} to $1.7 \times 10^{-4} \text{ ft}^{-1}$ (Johnson and others, 1968) (table 3). At this site, the larger value was measured in a sample from the Corcoran Clay (Johnson and others, 1968) (table 3), which would be expected to be highly compressible, yet this value is smaller than Helm's (1978) model-derived value by a factor of about 1.3. This discrepancy may result from the scale difference: laboratory consolidation tests measure a small sample of an aquifer system, whereas calibrated models focus on larger thicknesses of aquifer systems. Furthermore, laboratory consolidation tests are done with the premise that the sample is undisturbed, which is nearly impossible, and do not mimic natural stresses on, or stress history of, the sample.

SUMMARY

This report summarizes the hydraulic and mechanical properties affecting ground-water flow and aquifer-system compaction in the San Joaquin Valley, California. Because most storage values presented are components of the total aquifer-system storage and include inelastic and elastic skeletal storage values for aquifers and aquitards, the relations of components of aquifer-system storage to total aquifer-system storage were reviewed. Vertical hydraulic conductivity values generally are for discrete thicknesses of sediments—usually aquitards. The property values were obtained from publications that report the values as results of aquifer tests, stress-strain analyses of borehole extensometer observations, laboratory consolidation tests, and model simulations. These values will be used by the USBR during the calibration process of the WESTSIM model, which will simulate ground-water flow and land subsidence in the western San Joaquin Valley.

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