Prepared in cooperation with the Commonwealth of Massachusetts
Office of the State Geologist and Executive Office of Environmental Affairs

Surficial Geologic Map of the Salem Depot-Newburyport East-Wilmington-Rockport 16-Quadrangle Area in Northeast Massachusetts

Compiled by Byron D. Stone, Janet Radway Stone, and Mary L. DiGiacomo-Cohen

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Cover figure. View of bouldery end moraine deposits in Dogtown Commons, Gloucester, Massachusetts.

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Surficial Geologic Map of the Salem Depot-Newburyport East-Wilmington-Rockport 16-Quadrangle Area in Northeast Massachusetts

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Introduction

The surficial geologic map of the Salem Depot-Newburyport East-Wilmington-Rockport area shows the distribution of nonlithified earth materials at land surface in sixteen 7.5-minute quadrangles (total 658 mi^2) in northeast Massachusetts (fig. 1). Across Massachusetts, these materials range from a few feet to more than 500 ft in thickness. They overlie bedrock, which crops out in upland hills and in resistant ledges in valley areas. The geologic map differentiates surficial materials of Quaternary age on the basis of their lithologic characteristics (grain size, sedimentary structures, mineral and rock-particle composition), constructional geomorphic features, stratigraphic relationships, and age. Surficial materials also are known in engineering classifications as unconsolidated soils, which include coarse-grained soils, fine-grained soils, or organic fine-grained soils. Surficial materials underlie and are the parent materials of modern pedogenic soils, which have developed in them at the land surface. Surficial earth materials significantly affect human use of the land and an accurate description of their distribution is particularly important for water resources, construction aggregate resources, earth-surface hazards assessments, and land-use decisions.

The mapped distribution of surficial materials that lie between the land surface and the bedrock surface is based on detailed geologic mapping of 7.5-minute topographic quadrangles, as part of a cooperative state-wide mapping program between the U.S. Geological Survey and the Massachusetts Department of Public Works (now Massachusetts Highway Department) (Page, 1967, Stone, 1982), and the Office of the Massachusetts State Geologist. Each published geologic map presents a detailed description of local geologic map units, the genesis of the deposits, and age correlations among units. Regional summaries of these maps and unpublished maps discuss the ages of multiple glaciations, the nature of glaciofluvial, glaciolacustrine, and glaciomarine deposits, and the processes of ice advance and retreat across Massachusetts (Warren and Stone, 1986, Koteff and Pessl, 1981, papers in Larson and Stone, 1982, Oldale and Barlow, 1986, Stone and Borns, 1986).

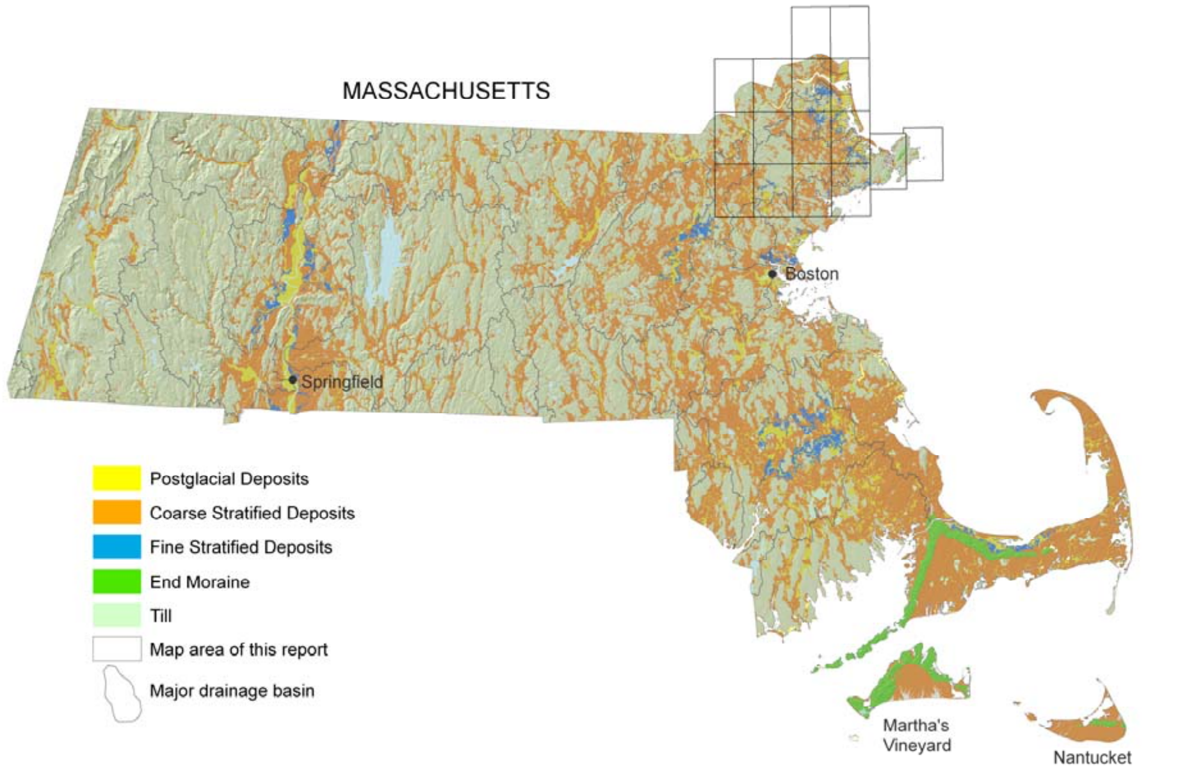


Figure 1. General distribution of glacial and postglacial deposits in Massachusetts (Stone and Beinikis, 1993, MassGIS, 1999) and map area covered by this report.

This compilation of surficial geologic materials is an interim product that defines the areas of exposed bedrock, and the boundaries between glacial till, glacial stratified deposits, and overlying postglacial deposits. This work is part of a comprehensive study to produce a statewide digital map of the surficial geology at a 1:24,000-scale level of accuracy. This map of 16 quadrangles revises previous digital surficial geologic maps (Stone and Beinikis, 1993, MassGIS, 1999) that were compiled on base maps at regional scales of 1:250,000 and 1:125,000. The purpose of this study is to provide fundamental geologic data for the evaluation of natural resources, hazards, and land information within the Commonwealth of Massachusetts.

Surficial Materials in Massachusetts

Most of the surficial materials in Massachusetts are deposits of the last two continental ice sheets that covered all of New England in the latter part of the Pleistocene ice age (Schafer and Hartshorn, 1965; Stone and Borns, 1986; Oldale and others, 1982). The glacial deposits are divided into two broad categories, ***glacial till*** and ***glacial stratified deposits***. Till, the most widespread glacial deposit, was laid down directly by glacier ice. Glacial stratified deposits are concentrated in valleys and lowland areas and were laid down by glacial meltwater in streams, lakes, and the sea in front of the retreating ice margin during the last deglaciation. Postglacial sediments, primarily floodplain alluvium and swamp deposits, make up a lesser proportion of the unconsolidated materials.

Glacial till deposits consist of nonsorted, generally nonstratified mixtures of mineral and rock particles, ranging in grain size from clay to large boulders. The matrix of most tills is composed dominantly of fine sand and silt. Boulders within and on the surface of tills range from sparse to abundant. Some tills contain lenses of sorted sand and gravel and less commonly, masses of laminated fine-grained sediments. The color and lithologic characteristics of till deposits vary across Massachusetts, but generally reflect the composition of the local underlying and northerly adjacent bedrock from which the till was derived. Till blankets the bedrock surface in variable thickness, ranging from a few inches to more than 200 ft, and commonly underlies stratified meltwater deposits. Tills deposited during the last two glaciations occur in superposition within Massachusetts (Koteff, 1966; Newton, 1978; Weddle and others, 1989). The upper till was deposited during the last (late Wisconsinan) glaciation; it is the most extensive till and commonly is observed in surface exposures, especially in areas where till thickness is less than 15 ft (thin till unit on the map). The lower till ("old" till) was deposited during an earlier glaciation (probably Illinoian), and has a more limited distribution; it is principally a subsurface deposit that constitutes the bulk of material in drumlins and other hills where till thickness is greater than 15 ft. The distribution of lower till is shown primarily by the thick till unit on the map. Generally, the lower till generally is overlain by thin upper till deposits in these areas. In all exposures showing the superposed two tills, the base of the upper till truncates the weathered surface of the lower till. The lower part of the upper till commonly displays a zone of shearing, dislocation, and brecciation, in which clasts of lower till were mixed and incorporated into the upper till during the last glaciation.

End moraine deposits are composed predominantly of bouldery ablation till, but may also locally include sorted sediments. These deposits were laid down by glacial-melting processes along active ice margins during retreat of the last (late Wisconsinan) ice sheet. Extensive end moraines on Nantucket and Martha’s Vineyard (fig. 1) are related to the terminal position of the late Wisconsinan ice sheet, and the end moraines on Cape Cod are associated with recessional positions of the last ice sheet. Less extensive end moraines occur locally elsewhere in southeastern Massachusetts, in the Boston area, and on this map in the Gloucester-Rockport area of northeastern Massachusetts.

Glacial stratified deposits consist of layers of well-sorted to poorly sorted gravel, sand, silt, and clay laid down by flowing meltwater in glacial streams, lakes, and marine embayments that occupied the valleys and lowlands of Massachusetts during retreat of the last ice sheet. Textural variations within the meltwater deposits occur both areally and vertically because meltwater-flow regimes were different in glaciofluvial (stream), glaciodeltaic (where a stream entered a lake or the sea), glaciolacustrine (lake bottom), and glaciomarine (marine bottom) depositional environments. Grain-size variations also resulted from meltwater deposition in positions either proximal to or distal from the retreating glacier margin, which was the principal sediment source. A common depositional setting contained a proximal, ice-marginal meltwater stream in which horizontally bedded glaciofluvial gravel and/or sand and gravel were laid down; farther down valley, the stream entered a glacial lake where glaciodeltaic sediments were deposited, consisting of horizontally layered sand and gravel delta-topset beds overlying inclined layers of sand in delta-foreset beds. Farther out in the glacial lake, very fine sand, silt, and clay settled out on the lake bottom in flat-lying, thinly bedded glaciolacustrine layers. Thick sequences having these textural variations commonly are present in the vertical section of meltwater deposits across the State (Stone and others, 1992). Detailed geologic maps permit precise mapping of meltwater sedimentary units within each glacial lake or valley outwash system (Jahns, 1941; 1953; Koteff, 1966). These units, known as *morphosequences* (Koteff, 1974; Koteff and Pessl, 1981), are the smallest mappable stratigraphic units on detailed geologic maps. *Morphosequences* are bodies of stratified meltwater sediments that are contained in a continuum of landforms, grading from ice-contact forms (eskers, kames) to non-ice-contact forms (flat valley terrace, delta plains) that were deposited simultaneously at and beyond the margin of the ice sheet and, graded to a specific base level. Each morphosequence consists of a proximal part (head) deposited within or near the ice margin and a distal part deposited farther away from the ice margin. Both grain size and ice-melt collapse deformation of beds decrease from the proximal to the distal part of each morphosequence. The head of each morphosequence is either ice marginal (ice contact) or near ice marginal. The surface altitude of fluvial sediments in each morphosequence was controlled by a specific base level, either a glacial-lake or marine water plane or a valley knickpoint. Few morphosequences extend distally more than 10 km, and most are less than 2 km in length. In any one basin, individual morphosequences were deposited sequentially as the ice margin retreated systematically northward. Consequently, in many places the distal, finer grained facies of a younger morphosequence stratigraphically overlies the proximal, coarse-grained facies of a preceding morphosequence. Figure 2 shows the variability of sediment types in the subsurface of glacial stratified deposits. The figure schematically shows the relationship between coarse-grained deltaic deposits and extensive fine-grained lake (or marine) deposits in the subsurface. Such coarse- and fine-grained units are common in most of the valleys and lowlands of Massachusetts (Langer, 1979, Stone and others, 1979; Stone and others, 1992; Stone and others, 2005). On this interim map, coarse-grained and fine-grained textural variations within glacial stratified deposits are shown only where they occur at land surface. Subsurface textural variations are not shown.

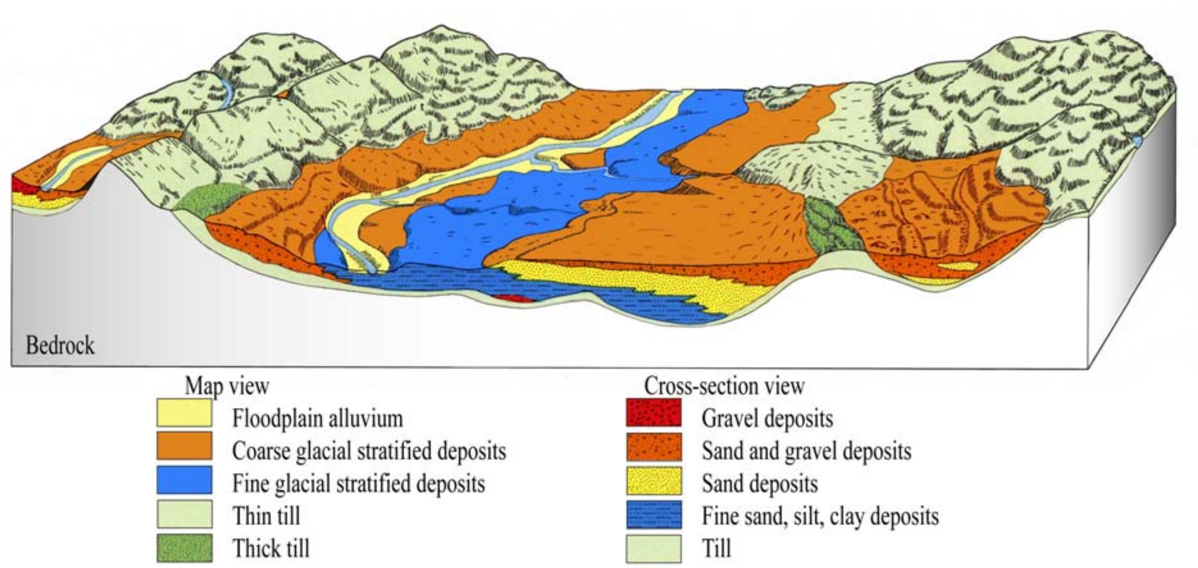


Figure 2. Block diagram illustrating the typical areal and vertical distribution of glacial and postglacial deposits overlying bedrock (modified from Stone and others, 1992).

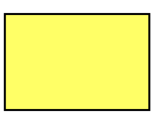
The areal distribution of till and stratified deposits across Massachusetts is related to regional physiography (fig.1). The thickness of these materials varies considerably because of such factors as the high relief of the bedrock surface, changing environments of deposition during deglaciation, and various effects of postglacial erosion and removal of glacial sediments. In highland areas, notably in the western and central parts of the State, till is the major surficial material and is present as a discontinuous mantle of variable thickness over the bedrock surface. Till is thickest in drumlins (reportedly as much as 230 ft thick) and on the northwest slopes of most bedrock hills. Glacial meltwater deposits that average 50 feet in thickness (Stone and Beinikis, 1993) overlie the till in small upland valleys and north-sloping basins between bedrock hills. Glacial stratified deposits are the predominant surficial materials in the Connecticut River valley, the northeastern and southeastern lowlands, and on Cape Cod and the islands. In northeastern Massachusetts, glacial meltwater deposits include glaciomarine sediments deposited when sea level was as much as 100 ft above present level. Glacial meltwater deposits generally overlie till; however, well logs indicate that in some places till is not present and the stratified deposits lie directly on bedrock. On Cape Cod and the islands, in the southeastern lowland, and in parts of the Connecticut River valley these deposits completely cover the till-draped bedrock surface.

Postglacial deposits locally overlie the glacial deposits throughout the State. Alluvium underlies the floodplains of most streams and rivers. Swamps occur in low-lying, poorly drained areas in upland and lowland settings, but swamp deposits are shown only where they are estimated to be at least 3 ft thick. Salt-marsh and estuarine deposits are present mainly along the tidal portions of streams and rivers entering the offshore areas. Beach deposits occur along the shoreline. Early postglacial regressive deposits represent former, higher marine shoreline positions and are preserved at altitudes 20 ft to 60 ft above modern sea level between Gloucester and the New Hampshire state line.

Description of Map Units

Postglacial Deposits

Artificial fill—Earth materials and manmade materials that have been artificially emplaced, primarily in highway and railroad embankments, and in dams; may also include landfills, urban development areas, and filled coastal wetlands



Floodplain alluvium—Sand, gravel, silt, and some organic material, stratified and well sorted to poorly sorted, beneath the floodplains of modern streams. The texture of alluvium commonly varies over short distances both laterally and vertically, and generally is similar to the texture of adjacent glacial deposits. Along smaller streams, alluvium is commonly less than 5 ft thick. The most extensive deposit of alluvium on the map is along the Merrimack River where the texture is predominantly sand, fine gravel, and silt, and total thickness is as much as 25 ft. Alluvium typically overlies thicker glacial stratified deposits.

Swamp deposits—organic muck and peat that contain minor amounts of sand, silt, and clay, stratified and poorly sorted, in kettle depressions or poorly drained areas. Most swamp deposits are less than about 10 ft thick. Swamp deposits overlie glacial deposits or bedrock. They locally overlie glacial till even where they occur within thin glacial meltwater deposits.

Salt-Marsh deposits—peat and organic muck interbedded with sand and silt, deposited in environments of low wave energy along the coast and in river estuaries. Marsh deposits are dominantly peat and muck, generally a few feet to 25 ft thick. In the major estuaries, marsh deposits locally overlie estuarine deposits (not mapped) which are sand and silt with minor organic material as much as 30-80 ft thick. The marsh and estuarine deposits generally are underlain by adjacent glacial material, either till, coarse stratified deposits, or glaciomarine fine deposits.

Beach and Dune deposits—sand and fine gravel deposited along the shoreline by waves and currents, and by wind action. The texture of beach deposits varies over short distances and is generally controlled by the texture of nearby glacial materials exposed to wave action. Sand beach deposits are composed of moderately sorted very coarse to fine sand, commonly laminated. Coarser layers may contain some fine gravel particles; finer layers may contain some very fine sand and silt. Gravel beach deposits are composed of granule-to-cobble-sized clasts in moderately sorted thin beds. Minor amounts of sand within gravel beds, and thin beds of sand as alternating layers. Beach deposits are rarely more than a few feet thick. Sand dune deposits are composed of moderately to well sorted, fine to medium sand, variably massive, laminated, and cross bedded. Dune deposits are as much as 35 ft thick. Unit includes artificial sand deposits in locally replenished beaches.

Early postglacial marine regressive deposits—sand and gravel deposited along former, higher shorelines in northeastern Massachusetts by waves and currents, and by wind action, as well as sand and gravel deposited by fluvial and wave action in lower estuarine valleys. The fluvial estuarine-terrace deposits are shown on the map where they overlie glaciomarine fine deposits; elsewhere, sand and gravel in postglacial terrace deposits are included in the glacial coarse stratified map unit. The fluvial terrace deposits are mixtures of gravel and sand within individual layers, and as alternating layers. Sand and gravel layers generally range from 25 to 50 percent gravel particles and from 50 to 75 percent sand particles. Layers are well to poorly sorted. Beach and near-shore deposits are composed of moderately sorted very coarse to fine sand, commonly laminated. Coarser layers may contain some fine gravel particles; finer layers may contain some very fine sand and silt. Regressive beach and near-shore deposits are rarely more than a few feet thick. Regressive spit deposits are 10-30 ft thick.

Glacial Stratified Deposits

Sorted and stratified sediments composed of gravel, sand, silt, and clay (as defined in particle size diagram, fig. 3) deposited in layers by glacial meltwater. These sediments occur as four basic textural units—gravel deposits, sand and gravel deposits, sand deposits, and fine deposits. On this interim map, gravel, sand and gravel, and sand deposits are not differentiated and are shown as *Coarse Deposits* where they occur at land surface. *Fine Deposits* also are shown where they occur at land surface. **Textural changes occur both areally and vertically (fig. 2), however subsurface textural variations are not shown on this interim map.**

PARTICLE DIAMETER										
in.	10	2.5	0.16	0.08	0.04	0.02	0.01	0.005	0.0025	0.00015
mm	256	64	4	2	1	0.5	0.25	0.125	0.063	0.004

Boulders	Cobbles	Pebbles	Granules	Very coarse sand	Coarse sand	Medium sand	Fine sand	Very fine sand	Silt	Clay
GRAVEL PARTICLES				SAND PARTICLES				FINE PARTICLES		

Figure 3. Grain-size classification used in this report, modified from Wentworth (1922).

Coarse deposits include: *Gravel deposits* composed mainly of gravel-sized clasts; cobbles and boulders predominate; minor amounts of sand within gravel beds, and sand comprises few separate layers. Gravel layers generally are poorly sorted and bedding commonly is distorted and faulted due to postdepositional collapse related to melting of ice. *Sand and gravel deposits* composed of mixtures of gravel and sand within individual layers and as alternating layers. Sand and gravel layers generally range from 25 to 50 percent gravel particles and from 50 to 75 percent sand particles. Layers are well to poorly sorted; bedding may be distorted and faulted due to postdepositional collapse. *Sand deposits* composed mainly of very coarse to fine sand, commonly in well-sorted layers. Coarser layers may contain up to 25 percent gravel particles, generally granules and pebbles; finer layers may contain some very fine sand, silt, and clay.

Glaciolacustrine Fine deposits include very fine sand, silt, and clay that occurs as well-sorted, thin layers of alternating silt and clay, or thicker layers of very fine sand and silt deposited in glacial lakes. Very fine sand commonly occurs at the surface and grades downward into rhythmically bedded silt and clay varves. Locally, this map unit may include areas underlain by fine sand.

Glaciomarine fine deposits include silty clay, fine sand, and some fine gravel deposited in a higher level sea in environments of low wave energy along the coast and in river estuaries. Fine to very fine sand, massive and laminated, commonly is present at the surface and grades downward into interbedded very fine sand, silt, and silty clay. Lower silty clay and clay is massive and thinly laminated. Total thickness is generally a few feet to 75 ft.

Glacial Till Deposits

Thin till—Nonsorted, nonstratified matrix of sand, some silt, and little clay containing scattered gravel clasts and few large boulders; in areas where till is generally less than 10-15 ft thick and including areas of bedrock outcrop where till is absent. Predominantly upper till of the last glaciation; loose to moderately compact, generally sandy, commonly stony. Two facies are present in some places; a looser, coarser-grained ablation facies, melted out from supraglacial position; and an underlying more compact, finer-grained lodgement facies deposited subglacially. In general, both ablation and lodgement facies of upper till derived from fine-grained bedrock are finer grained, more compact, less stony and have fewer surface boulders than upper till derived from coarser grained crystalline rocks. Fine-grained bedrock sources include the red Mesozoic sedimentary rocks of the Connecticut River lowland, marble in the western river valleys, and fine-grained schists in upland areas.

Thick till—Nonsorted, nonstratified matrix of sand, some silt, and little clay containing scattered gravel clasts and few large boulders at the surface; in the shallow subsurface, compact, nonsorted matrix of silt, very fine sand, and some clay containing scattered small gravel clasts in areas where till is greater than 10-15 ft thick, chiefly in drumlin landforms in which till thickness commonly exceeds 100 ft (maximum recorded thickness is 230 ft). Although upper till is the surface deposit, the lower till constitutes the bulk of the material in these areas. Lower till is moderately to very compact, and is commonly finer grained and less stony than upper till. An oxidized zone, the lower part of a soil profile formed during a period of interglacial weathering, is generally present in the upper part of the lower till. This zone commonly shows closely spaced joints that are stained with iron and manganese oxides.

End moraine deposits – Composed predominantly of boulders and ablation facies sandy upper till; lenses of stratified sand and gravel occur locally within the till. Surface boulders on end moraine deposits are generally more numerous than on adjacent till surfaces; dense concentrations of boulders are present in some places. Deposits occur as free-standing hummocky landforms, commonly in ridges that trend NE – SW, and range in thickness from 10 to 60 ft.

Bedrock Areas

Bedrock outcrops and areas of abundant outcrop or shallow bedrock— Solid color shows extent of individual bedrock outcrops; line pattern indicates areas of shallow bedrock or areas where small outcrops are too numerous to map individually; in areas of shallow bedrock, surficial materials are less than 5-10 ft thick.

Map Compilation

This compilation shows surficial geology in sixteen 7.5-minute quadrangles in northeast Massachusetts: Exeter, Hampton, Salem Depot, Haverill, Newburyport West, Newburyport East, Lawrence, South Groveland, Georgetown, Ipswich, Wilmington, Reading, Salem, Marblehead North, Gloucester, and Rockport (fig. 4, fig 5 area B). It is the second in a series of interim products intended to cover the entire State of Massachusetts. Figure 5 shows all of the compilation areas for surficial geology in Massachusetts. These maps will be produced sequentially by letter designation.

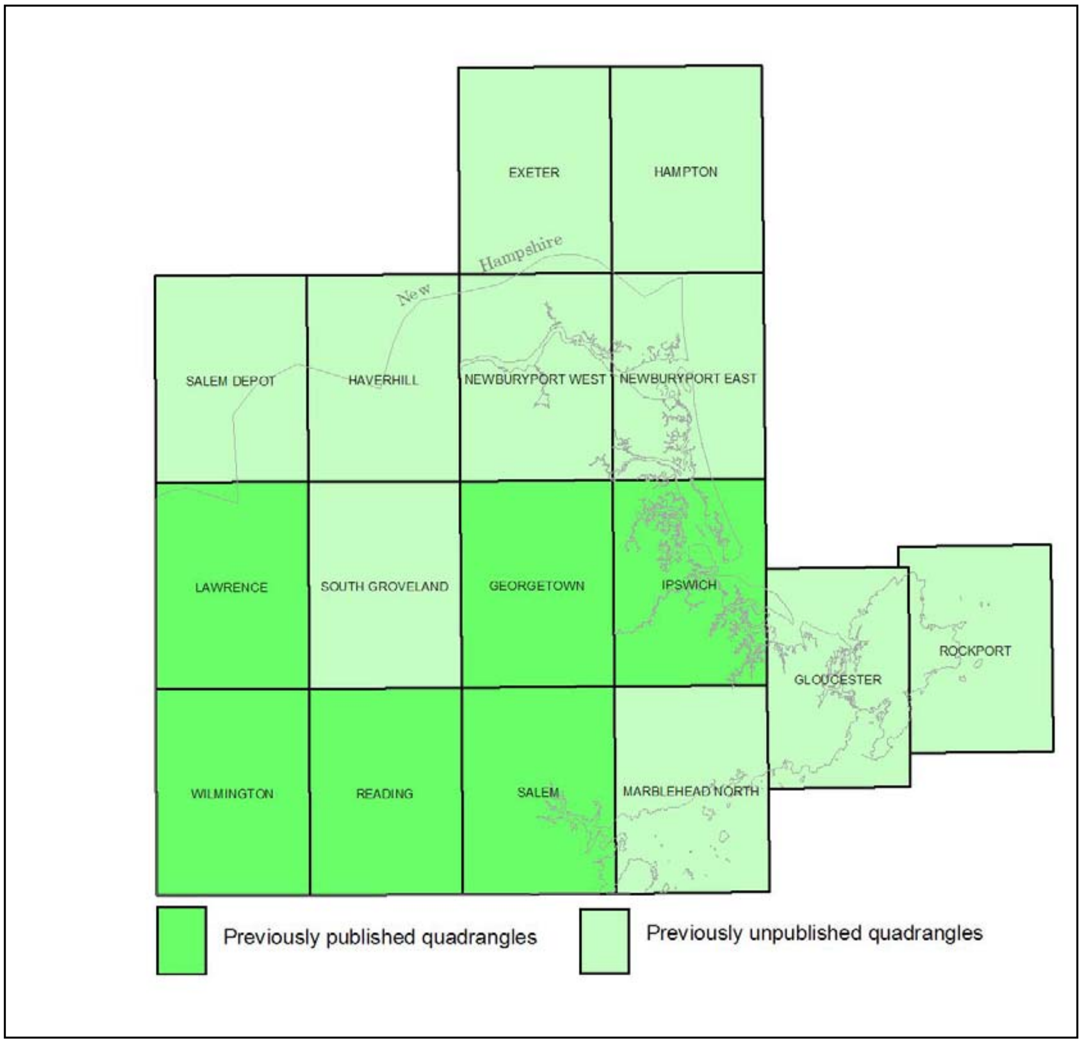


Figure 4. 7.5-minute quadrangles in this compilation.

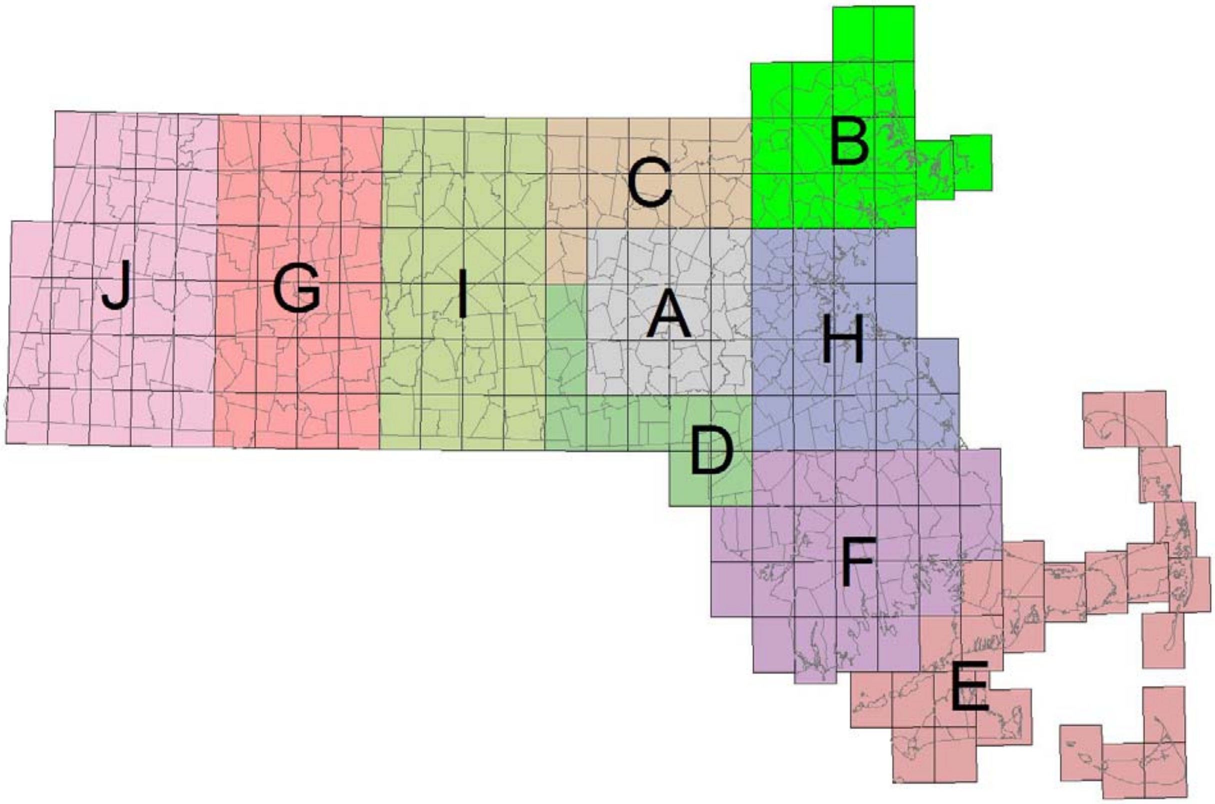


Figure 5. Compilation areas in Massachusetts.

This map was compiled in several steps: 1) Paper copies of the published surficial geologic maps for six quadrangles were scanned and georeferenced by MassGIS. 2) The Office of the Massachusetts State Geologist vectorized the georeferenced images in order to digitally retain the original line work of the published maps (Mabee and others, 2004). 3) Digital geologic map units were compiled and grouped into thirteen basic units in four broader categories: *Postglacial deposits* including artificial fill, floodplain alluvium, swamp deposits, salt marsh deposits, beach and dune deposits, and early postglacial regressive deposits; *glacial stratified deposits* including coarse-grained and fine-grained deposits, *glacial till* including thin till, thick till (drumlins), and end moraine deposits: and *bedrock areas* including outcrops and areas of shallow bedrock). The distribution of glacial stratified deposits beneath adjacent overlying postglacial deposits and water bodies was inferred by the compilers. 4) The same basic units for ten unpublished quadrangles were compiled and digitized from scanned field maps by U.S. Geological Survey personnel. 5) The 16 individual quadrangles were joined and edge-matched in order to form a seamless geologic map. Discrepancies along quadrangle boundaries were resolved, and thick till areas were added by the compilers in quadrangles where this unit was not previously mapped.

All geologic mapping was completed at 1:24,000-scale; however the browse graphic is presented at 1:50,000-scale with shaded relief base. The 1:24,000-scale, 10-ft contour interval topographic base maps used for this mapping effort are included as part of the digital data package in the GIS\24k_basemaps folder. The GIS\shapes folder included with this report contains 4 ARCGIS shapefiles which are geologic units that cover the entire map area, and are intended for use at quadrangle scale; the shapefiles can be clipped by quadrangle or town boundaries. Unlike conventional geologic maps, the digital mapping is arranged in layers according to superposition. The till-bedrock shapefile should be placed on the bottom, and overlain by the stratified deposits shapefile; these materials are shown everywhere that they occur including beneath postglacial deposits, such as swamp deposits, floodplain alluvium, and water bodies. The early_postglacial and postglacial shapefiles should be placed on top because these materials overlie the other two layers. Instructions for using the digital files are included in the README file and metadata.

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Appendix

Sources of Data by 7.5-Minute Quadrangle

Exeter
Map units were reproduced from Goldsmith (2001)

Hampton
Map units were reproduced from Koteff, Carl, Gephart, G. D., and Schaefer, J.P. (1989)

Salem Depot
Stone, B.D., 1982, Unpublished field map

Haverhill
Stone, B.D., 1982, Unpublished field map

Newburyport West
Stone, B.D., 1982, Unpublished field map

Newburyport East
Stone, B.D., 1982, Unpublished field map

Lawrence
Map units were reproduced from Castle (1958); Glacial Stratified Deposits in this quadrangle include deposits of glacial Lake Methuen, and other glaciofluvial and glaciolacustrine deposits. Fine-grained glacial stratified deposits at land surface include lake-bottom deposits of glacial Lake Methuen (unit Qml of Castle, 1958); the fine-grained unit has been extended beneath adjacent water bodies and postglacial deposits on this map. Deposits mapped as morainal ice-contact deposits (unit Qmi of Castle, 1958) have been included here as coarse stratified deposits. Thick till areas were reproduced from the drumlin till unit (Qgd of Castle, 1958); other areas of thick till were inferred from photographic image and topographic analysis.

South Groveland
Stone, B.D., 1982, Unpublished field maps

Georgetown
Map units were reproduced from Cuppels (1969); Glacial Stratified Deposits in this quadrangle include glaciofluvial, glaciolacustrine, and glaciomarine deposits. Fine-grained glacial stratified deposits at land surface include parts of marine and estuarine deposits (unit Qm of Cuppels, 1969); the fine-grained unit has been extended beneath adjacent water bodies and postglacial deposits on this map. Thick till areas shown on this map were inferred from photographic image and topographic analysis and drumlin symbols shown by Cuppels (1969).

Ipswich
Map units were reproduced from Sammel (1963); Glacial Stratified Deposits in this quadrangle include glaciofluvial, glaciolacustrine, and glaciomarine deposits. Fine-grained glacial stratified deposits at land surface include marine and estuarine clay deposits (unit Qmc of Sammel, 1963); this unit has been extended beneath adjacent water bodies and postglacial deposits on this map. Thick till areas were reproduced from the drumlin till unit (Qd of Sammel, 1963); other areas of thick till were inferred from photographic image and topographic analysis.

Wilmington
Map units were reproduced from Castle (1959); Glacial Stratified Deposits in this quadrangle include glaciofluvial and glaciolacustrine deposits. Thick till areas were reproduced from the drumlin till unit (Qgd of Castle, 1959); other areas of thick till were inferred from photographic image and topographic analysis.

Reading
Map units were reproduced from Oldale (1962); Glacial Stratified Deposits in this quadrangle include glaciofluvial and glaciolacustrine deposits. Fine-grained glacial stratified deposits at land surface include glacial lake deposits (unit Ql of Oldale, 1962); the fine-grained unit has been extended beneath adjacent water bodies and postglacial deposits on this map. Thick till areas were reproduced from the drumlin till unit (Qgd of Oldale, 1962) and other areas of thick till were inferred form photographic image and topographic analysis\.

Salem
Map units were reproduced from Oldale (1964); Glacial Stratified Deposits in this quadrangle include glaciofluvial, glaciolacustrine, and glaciomarine deposits. Fine-grained glacial stratified deposits at land surface include marine clay deposits (unit Qmd of Oldale, 1964); the fine-grained unit has been extended beneath adjacent water bodies and postglacial deposits on this map.

Marblehead North
Map units were reproduced from Carnevale, 1979.

Gloucester
Stone, B.D., 1982, Unpublished field maps

Rockport
Stone, B.D., 1982, Unpublished field maps