Volcanic deformation models for Deception Island (South Shetland Islands, Antarctica)

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Summary  GPS observations from a Deception Island Volcano geodetic network (South Shetland Islands, Antarctica) during 1991/1992 to 2002/2003 campaigns have characterized the deformation of the island resulting from its volcanic activity. From January 1992 to December 1999, radial extension together with an uplift episode was detected. During this period two important episodes of seismic activity took place, in particular, during the campaigns 1991/92 and 1998/99. From 2000 until 2003, no large displacements were detected, although the stations seem to converge in a radially compressive manner together with a subsidence, reflecting a change in the geodynamics of the island. Two clear alignments are observed on the contour maps that characterize the deformation, one in the NW-SE direction until 2000 and another in the NE-SW direction from 2000 to 2003, coinciding with the principal directions of the tectonic features of the region, the Hero Fracture Zone and the Bransfield Rift.


Introduction

Deception Island is located in a very complex tectonic zone where several tectonic plates converge: The South American Plate, the Antarctic Plate and three minor plates, Scotia micro-plate, Phoenix micro-plate and South Shetland micro-plate (Baraldo, 1999). The former Phoenix microplate is now part of the Antarctic plate. It is delineated by the former spreading center (Drake Ridge) to the northwest, the Hero Fracture Zone to the southwest, the South Shetland subduction system to the southeast and the Shackleton Fracture Zone to the northeast. The microplate was formerly being subducted under the South Shetland Islands and Bransfield Strait, but several studies reveal that the subduction process is no longer active.

The active volcanism in this area is mainly located in the extensional core of Bransfield rift and is characterized by the presence of several emerged volcanoes like Deception, Penguin and Bridgeman Islands. In addition there are also numerous submarine volcanic centres which show different evolution stages, from the volcanic cone to its collapse, generated by the existing divergent tectonics (Canals, et al., 1997). These volcanic centres are 15 km base diameter and about 450 m high. In addition, there are around 34 small centres with a mean diameter of 2.5 km. These centres are aligned following the main direction of the basin.

The continuous volcanic activity in this area is manifested by historical eruptions at Deception and Penguin islands, together with the Bransfield central basin rifting and the magmatic processes in the rift during the last two million years. This fact suggests that the volcanic activity in the Bransfield basin is in continuous development (González-Ferrán, 1991).

REGID geodetic network

The geodetic network for Deception Island (REGID) was established from observations of GPS System satellites in order to set up the basic reference frame not only for the geodetic studies but also for other research. In this way, the network is the fundamental reference for the Spanish cartography in Antarctica, for geophysical and oceanographic observations and for hydrographical sounding carried out throughout the Spanish campaigns in the Antarctic. On the other hand, the GPS System accuracy, especially for horizontal positioning, makes the REGID network very useful for understanding geodynamic activity of the island.

Stations of the network

The establishment of the geodetic network began during the 1989-90 Antarctic campaign. The aim of the studies and the place where they were going to be carried out led to the construction of high stability benchmarks, with no visual impact. They were built with concrete, well-rooted in the permafrost with steel bars and at a low height above the ground. A standard screw was fixed in one of the corners of the post.
for the use of geodetic instrumentation. Furthermore, some of the stations were built using existing structures of the buildings destroyed by the volcanic activity which happened between 1967 and 1970 (Berrocoso, 1997).

In the 1990-91 campaign, the geodetic network consisted of five stations, four of them in Deception Island (BARG, FUMA, PEND, BALL) and another one at the Spanish Base Juan Carlos I (BEJC) in Livingston Island. Their placements were chosen according to several factors, including accessibility, stability of the building, proximity of fumarolic areas, density of located epicenters, last eruptions. During the 1995-96 campaign a new station was built near the Spanish Antarctic Base Gabriel de Castilla. This station (BEGC) was set as the new main point of the network. It has electric supply and it was built in the same way as the others (Berrocoso, 1997).

During the 2001-02 campaign seven more stations were built in the northern and southern part of Deception Island, in order to cover its entire inner ring and to make the REGID network coverage more consistent. The distribution of the stations is given in Figure 1. The construction of these new stations was carried out according to the same geodetic directions as the earlier ones: clear GPS horizon, low multipath effect. The area where they were built was chosen taking into account the volcanic nature of the island, but without considering any prior volcanic activity (Vila, 1992; García et al., 1997; Ibáñez et al., 2000).

**GPS data collection**

Stations of the initial network have been observed from 1991-92 during the Spanish Antarctic campaigns. BEGC station was included in the 1995-96 campaign, when it was observed for the first time. After the construction of seven new stations in the 2001-02 campaign, COLA, GEODEC, UCA1, TELE BOMB, CR70 and GLAN, the whole network has been observed during the austral summers periods since the 2001-02 campaign to present. The triangle determined by BEGC, FUMA and PEND is continuously observed during the whole surveying from 2002-03 campaign. Antennae height corrections and phase centre offsets were applied to the data. The sampling rate was set to 15 s. and the elevation mask was 10º.

**Data processing and adjustment of the network**

Data from GPS surveying in Deception Island were processed using the Bernese v4.2 software (Beutler et al., 2001) using precise ephemerides. Data from the IGS station OHIG was used to the adjustment of the network. It is situated at the Chilean Base O’Higgins in the Antarctic Peninsula, 150 km. away from Deception Island.

Every campaign was processed independently, fixing the coordinates of one of the stations of the network. BARG was the fixed station until 2001, and then it was changed to BEGC station. In the 1991-92 campaign, the coordinates for the fixed station BARG were obtained from the international campaign SCAR92. The precise absolute coordinates for the station were referred to the ITRF91, epoch 92.2. To get these coordinates for the next campaigns, in 1995-96 and 2001-02, the coordinates of OHIG corresponding to ITRF96, epoch 96.1 and to ITRF97, epoch 99.9, were fixed.

### REGID WGS84 COORDINATES

<table>
<thead>
<tr>
<th>STATION</th>
<th>Φ (S)</th>
<th>λ (W)</th>
<th>h (m)</th>
<th>σ_φ(m)</th>
<th>σ_λ(m)</th>
<th>σ_h(m)</th>
</tr>
</thead>
<tbody>
<tr>
<td>OHIG</td>
<td>63.19 15.8930</td>
<td>57.54 4.7999</td>
<td>32.469</td>
<td>FIXED</td>
<td>FIXED</td>
<td>FIXED</td>
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<tr>
<td>BEGC</td>
<td>62.58 43.6576</td>
<td>60.40 27.5304</td>
<td>82.070</td>
<td>0.001</td>
<td>0.001</td>
<td>0.002</td>
</tr>
<tr>
<td>BARG</td>
<td>62.58 30.2679</td>
<td>60.41 53.3601</td>
<td>22.283</td>
<td>0.001</td>
<td>0.001</td>
<td>0.006</td>
</tr>
<tr>
<td>BALL</td>
<td>62.58 38.5553</td>
<td>60.33 52.5089</td>
<td>25.968</td>
<td>0.003</td>
<td>0.001</td>
<td>0.006</td>
</tr>
<tr>
<td>FUMA</td>
<td>62.57 41.0170</td>
<td>60.42 59.3471</td>
<td>22.911</td>
<td>0.001</td>
<td>0.001</td>
<td>0.006</td>
</tr>
<tr>
<td>BEJC</td>
<td>62.39 46.7782</td>
<td>60.23 19.9915</td>
<td>31.241</td>
<td>0.002</td>
<td>0.001</td>
<td>0.009</td>
</tr>
<tr>
<td>PEND</td>
<td>62.56 9.8456</td>
<td>60.35 34.3437</td>
<td>28.841</td>
<td>0.001</td>
<td>0.001</td>
<td>0.003</td>
</tr>
<tr>
<td>COLA</td>
<td>62.59 27.9180</td>
<td>60.37 31.8014</td>
<td>48.050</td>
<td>0.003</td>
<td>0.001</td>
<td>0.009</td>
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<td>GLAN</td>
<td>62.57 58.3621</td>
<td>60.35 23.8528</td>
<td>27.541</td>
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<td>0.003</td>
<td>0.012</td>
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<tr>
<td>GEOD</td>
<td>62.58 56.4127</td>
<td>60.39 11.7290</td>
<td>42.182</td>
<td>0.003</td>
<td>0.001</td>
<td>0.009</td>
</tr>
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<td>UCA1</td>
<td>62.56 28.4103</td>
<td>60.41 28.0891</td>
<td>28.667</td>
<td>0.003</td>
<td>0.001</td>
<td>0.009</td>
</tr>
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<td>CR70</td>
<td>62.55 23.6706</td>
<td>60.38 1.0013</td>
<td>23.600</td>
<td>0.003</td>
<td>0.003</td>
<td>0.015</td>
</tr>
<tr>
<td>TELE</td>
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<td>60.41 25.5485</td>
<td>23.791</td>
<td>0.003</td>
<td>0.003</td>
<td>0.012</td>
</tr>
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<td>60.39 33.8481</td>
<td>23.785</td>
<td>0.003</td>
<td>0.003</td>
<td>0.015</td>
</tr>
</tbody>
</table>

Table 1. REGID geodetic network stations WGS84 coordinates (based on ITRF2000.0).
Due to the proximity of the BEGC station to the Spanish Base Gabriel de Castilla, the receiver is easily maintained and it is possible to collect data continuously during the campaign. That is why this station was considered as the main one in the network REGID from the 2001-02 campaign. Its coordinates were obtained by fixing the ones of OHIG corresponding to the ITRF2000, epoch 2000.0.

Once the coordinates of the fixed station were set, the coordinates for the rest of the stations in the network were obtained by means of radial baselines from the fixed site to the rest of the stations. Simultaneous observations were used for the processing of the data. In the 1991-92 campaign, the radial processed baselines were BARG-FUMA, BARG-PEND and BARG-BALL. The ambiguities were calculated by the SIGMA strategy of the BERNESE software, since the observations were made by single-frequency receivers. Finally, the results were reference to the ITRF91; in 1995-96, observations of both L1 and L2 frequencies were available, so the ambiguities resolution was solved applying the QIF strategy. The final solution was referenced to the ITRF96; in the 1999-00, the baselines were processed altogether since it was the first time that every station was observed simultaneously. The ambiguities were also resolved with the QIF strategy and the final solution was referenced to the ITRF97; from 2001-02 campaign, solutions for every week were combined with the ADDNEQ program of the BERNESE software. QIF strategy was also used and the final solution was referenced to the ITRF2000. To estimate the tropospheric delays we use the default meteorological data, since the only available data were the medium temperature and pressure and their inclusion could produce more errors than the models provided by the software. The final results are shown in Table 1. They are referenced to the ITR2000, epoch 2000.0.

Volcanic deformation models determination

Deformation models representing Deception Island volcanic activity were obtained from the GPS observations of the stations in the REGID geodetic network. GPS observations have been episodically made from 1991-92 campaign to present. Taking into account that GPS system provides accuracy twice better in horizontal positioning than in the elevation component, we distinguish between the horizontal and the vertical models, although it is essential to use both results together for a proper interpretation. (Donellan, A., 1993; El-Fiky et al. 1999; Calais et al., 2000; Mantovani et al., 2001; Murray and Wooller, 2002).

![Figure 2. From left to right, up to bottom: horizontal and vertical deformation models for 1996.2/1999.9, and 2002.2/2003.2 epochs (cms/year), respectively.](image-url)
To obtain the deformation models a topocentric system was established with origin at BARG. In order to obtain the absolute deformation of the island, BEJC station on Livingston Island was considered. Displacement models were calculated comparing absolute topocentric coordinates between two consecutives campaigns (Figure 2); the velocity fields and its gradient were determined by a finite elements interpolation; horizontal deformation tensors were calculated to obtain the deformation parameters, that is, the extension-compression area, determining the components in NS and EW directions (Dermanis, 1985; Eren, 1984; Vaníček et al., 1986). Finally, the geodetic maximum deformation is obtained (Grafarend and Voosoghi, 2003). Figure 3 shows the mathematical horizontal deformation models obtained from the displacement models by dilatation (Bibby, 1982; Ward, 1998; Drew and Sray, 1989; Dzu risin, 2006).


Acknowledgments. The realization of this work has been possible thanks to the financial support to the following projects: ANT1999-1430-e/HESP, DECVOL, REN 2000-0551-C03-01 GEODEC, REN2000-2690-E Acquisition of a Scientific Software for GPS data processing, and CGL2004-21547-E/ANT, VOLTEDEC, all of them funded by the Spanish Ministry of Science and Technology through the National Program of Antarctic Research of Natural Resources.
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