

Opening of the Drake Passage: Does this event correlate to climate change and biotic events from the Eocene La Meseta Formation, Seymour Island, Antarctic Peninsula?

J. A. Case

College of Science, Health & Engineering, Eastern Washington University, Communications Bldg. 138, Cheney, WA 99004 (jcase@mail.ewu.edu)

Summary The time frame for opening of the Drake Passage, which resulted in the onset of Antarctic climatic cooling and then to the development of ice sheets on the Antarctic Peninsula, is hypothesized to be an early Oligocene event. Rock units from the topmost levels of the La Meseta Formation on Seymour Island, Antarctic Peninsula exhibit evidence of ice sheet formation. The date for ice sheet development is at the Eocene-Oligocene boundary. Thus the opening of the Drake Passage is hypothesized to be at Eocene-Oligocene boundary. However, fish teeth extracted from deep-sea cores were analyzed to provide data on a deepwater opening of the Drake Passage correlated to the Late Eocene (ca. 41 Ma). The data from vertebrate paleofaunas and associated paleofloras from the La Meseta Formation can be used to relate the opening of the Drake Passage to climatic indicators from these fossil remains.

Citation: Case, J.A. (2007), Opening of the Drake Passage: does this event correlate to climate change and biotic events from the Eocene La Meseta Formation, Seymour Island, Antarctic Peninsula? in Antarctica: A Keystone in a Changing World – Online Proceedings of the 10th ISAES X, edited by A. K. Cooper and C. R. Raymond et al., USGS Open-File Report 2007-1047, Extended Abstract 117, 3 p.

Introduction

The time frame for opening of the Drake Passage, which resulted in the development of the Antarctic Circumpolar Current, subsequent onset of Antarctic climatic cooling, and leading to the development of an earliest Oligocene ice sheet on the Antarctic Peninsula is considered to be an Oligocene event (Zachos et al., 2001). Ivany et al. (2006) described rock units from the topmost levels of the La Meseta Formation on Seymour Island, northeastern Antarctic Peninsula and overlying horizons, which indicate ice sheet formation. Data from dinoflagellates and strontium dating place the date for ice sheet development at the Eocene-Oligocene boundary. Thus the date for the opening of the Drake Passage is hypothesized to be at the Eocene-Oligocene boundary and correlated to a sharp drop in ocean temperature.

Scher and Martin (2006) utilized a rare earth element, neodymium (Nd), contained in fish teeth extracted from deep-sea cores to provide data on a deepwater opening of the Drake Passage. They used ratios of $^{143}\text{Nd}/^{144}\text{Nd}$ to determine the transition from non-radiogenic to radiogenic Nd values in Atlantic waters, which would mark the flow of the more radiogenic Pacific waters into the Atlantic ocean as the signal for the deepwater opening of the Drake Passage in the Late Eocene (ca. 41 Ma). A shallow water opening of the Drake Passage of less than 1000 meters deep may have formed as early as 50 Mya based on tectonic evidence in the Weddell Sea (Livermore et al., 2005). The formation of new tectonic basins in the region, that later formed the Scotia Arc and the northern Antarctic Peninsula, is correlated with a drop in Southern Ocean temperatures. Based on oxygen isotope data from benthic foraminifera, this is hypothesized to represent the initial shallow water opening between South America and Antarctica (Livermore et al., 2005).

La Meseta Formation

It is possible to examine these hypotheses regarding the time frame for the opening of the Drake Passage from what is currently known in a locality much closer to the site of the actual opening at Seymour Island (Isla Marambio). The La Meseta Formation (LMF) occupies nearly all of the northern third of the island and represents a 720 m thick, stratified, sequence of fossiliferous sands, silts, mudstones and shell beds, which spans nearly all (if not all) of the Eocene including the Eocene-Oligocene boundary (Ivaney et al., 2006). The oldest dates (52-54 Ma), based on $^{87}\text{Sr}/^{86}\text{Sr}$ isotopic ratios (Dutton et al., 2002) come from the 150 m level in Telm 2 of Sadler (1988) at the boundary of the Valle de las Focas and Acantilados Allomembers of Marenssi and Santillana (1994). Consequently, the base of the LMF should be close to the beginning of the Eocene at 54.8 Ma (Gradstein et al., 2004). Ivaney et al. (2006) report a date from the top of the LMF of 33.57-34.78 Ma ($^{87}\text{Sr}/^{86}\text{Sr}$) and this date is consistent with a previous date for the topmost LMF of 34.2 Ma (Dutton et al., 2002). Thus, the La Meseta Formation should contain a near complete record of the transition from the warm Early Eocene conditions, through the transitional cooling of the Middle Eocene to the greater cooling in the latest Eocene and finally to the onset of Early Oligocene ice sheet development. A question that can be asked is to what degree is there a correlation between the opening of the Drake Passage, ocean temperatures and climatic indicators based on megafossil evidence of the local paleofloras and the La Meseta paleofaunas?

With the potential of added dates within the LMF from Marenssi (2006) where he correlated the unconformities between the La Meseta stratigraphic units to low sea level stands of Haq et al. (1987), it is then possible to fit the Eocene portion of the ocean temperature curve of Zachos et al. (2001) to the La Meseta stratigraphy. Therefore, it is

possible to relate the ocean temperature fluctuations with the deep-water opening of the Drake Passage and the biotic events documented by the paleofloras and paleofaunas from the La Meseta Formation.

Paleofloras and paleofaunas

The Late Paleocene Nordenskjöld (= Cross Valley; Francis et al., 2003) paleoflora from the Cross Valley Formation on Seymour Island is dominated by angiosperm taxa with some thirty-six species, thirteen of which have entire leaf margins, while twenty-three species exhibit toothed leaf margins indicating a warm temperate forest (Francis et al., 2003). Additionally, there are both araucarian and podocarp conifers and at least three fern species (Case, 1988). When the leaf morphology of the angiosperm species are examined via CLAMP analysis, a warm Mean Annual Temperature (MAT) of 13.5°C is predicted for the Antarctic Peninsula with an average precipitation of 2110 mm/yr in the late Paleocene (Francis et al., 2003).

There is a distinct change between the Late Paleocene, Nordenskjöld paleoflora and the Early Eocene, Acantilados (= Cross Valley; Case, 1988) paleoflora from the Acantilados allomember. An increase in leaf size was noted in response to the warm Early Eocene conditions with presumably increased precipitation as well (Case, 1988).

The Middle Eocene, La Meseta paleoflora from the Cucullaea I allomember, exhibits a decrease in leaf size from the paleofloras both from the Nordenskjöld and the Acantilados paleofloras, suggesting a drop in temperature (Case, 1988). CLAMP analysis of the La Meseta paleoflora suggests only a small drop in MAT to 11-13°C but with a significant change to a highly seasonal climate with a Mean Winter Temperature (MWT) of -3 to 2°C (Gandolfo et al., 1998a,b). This cool temperate rainforest includes podocarp, araucarian and cupressacean gymnosperms and at least three species of fern (Case, 1988). Torres et al. (1994) indicate the fossil wood taxa have affinities with extant taxa living today in cold-temperate rainforests of the Valdivian region of southern South America (Poole and Cantrill, 2006). Modern counterparts of the Eocene mammal paleofauna from the La Meseta Formation, also live in these same Valdivian rainforests of southern Chile and Argentina (Woodburne and Case, 1996). This Valdivian rainforest-like environment represented by the La Meseta paleoflora with cold wet winters, but drier than that predicted for earlier paleofloras where precipitation was reduced to 1534 mm/yr in the Middle Eocene (Francis et al., 2003).

The Middle Eocene Antarctic terrestrial vertebrate paleofauna from the La Meseta Formation exhibits a wide, but bimodal, range of body sizes from small marsupials (also a rodent-like gondwanathere) to large-sized ungulates along with a large-bodied sloth and cursorial birds including ratites and phororachoid birds. The paleofauna lacks medium-sized marsupials and notoungulate mammals and thus the fauna has a U-shaped, bimodal distribution of body sizes. This bimodal body size distribution pattern also appears in higher latitude living mammalian faunas as a response to cold winter temperatures at these latitudes (Case, 2006). If the Patagonian Late Paleocene to Early Eocene mammalian paleofaunas are the source for the La Meseta Paleofauna (LMP), then a high degree of faunal similarity would be expected, but the LMP shows a distinct endemism at the generic or species level. The degree of endemism in the LMP may be the result of a vicariant barrier that could either be climatic (dropping of temperature, especially MWT) or physical (the opening of the Drake Passage) or the interaction of both vicariant phenomena (Reguero and Marennissi, in press).

The Marine Vertebrate Paleofauna from the La Meseta Formation shows some relationship with the changes in ocean temperature and potentially to the deep-water opening of the Drake Passage. Sharks are very abundant in Telm 3-5 (Acantilados through Cucullaea I (Case, 1992) with penguin fossils found frequently and the occasional whale material recovered in these same units. The vertebrate record changes significantly in Telm 7 (Submeseta) in that only the large ungulates and giant birds are present (Reguero and Marennissi, in press), while the smaller marsupials appear to drop out of the record. Among marine vertebrates, sharks have nearly disappeared, while the abundance of penguins dramatically increases and whale material is much more frequent.

The change in the vertebrate record in the mid- to upper portions of the La Meseta Formation are suggestive of a significant change and would certainly correlate well with a deep water opening of the Drake Passage at 41 Ma (Scher and Martin, 2006). The change from Middle Eocene Climatic Optimum to cooler Late Eocene conditions is consistent with the vertebrate faunal changes, which probably occurred in Telm 6 (Cucullaea II).

The point of uncertainty is whether the floral changes between the Early Eocene, Acantilados paleoflora and the Middle Eocene, La Meseta paleoflora and the proposed climatic changes are correlated with the originally shallow breaching of the Patagonian-Antarctic Isthmus resulting in the Drake Passage. Alternatively, these biotic changes could be the result of cooling caused by significant drops in atmospheric carbon dioxide levels (DeConto and Pollard, 2003), which would also have the potential to cause the development of ice sheets in the Antarctic Peninsula.

A full test and determination of the relationship of the opening of the Drake Passage to climatic events can only be achieved with further study, which is the intent of a current proposal by the author to the National Science Foundation.

Summary

The La Meseta Formation contains a very complete record of the Eocene timeframe in Antarctica. The effect of the opening of the Drake Passage (hypothesized to be at or prior to 41 Ma) on climatic and biotic change can be examined

through the paleofloras and vertebrate paleofaunas collected from that Formation. The degree of correlation between timing of the formation of the Drake Passage and climatic and biotic changes can then be tested.

Acknowledgements. I would like to thank Vanessa Thorn for her editorial assistance on this extended abstract. I would like to thank Marcelo Reguero of the Museo de La Plata and Sergio Marenni of the Instituto Antartico Argentino for their conversations and writings on the La Meseta Formation. I would also like to thank Jane Francis for inviting me to participate in the symposium and to the National Science Foundation, Office of Polar Programs for funding our Antarctic Research.

References

- Case, J. A. (1988), Paleogene floras from Seymour Island, Antarctic Peninsula, in *Geology and Paleontology of Seymour Island, Antarctic Peninsula*, edited by R. M. Feldmann and M. O. Woodburne, pp. 523-530, *Geol. Soc. Am. Mem.*, 169.
- Case, J. A. (1992), Evidence from fossil vertebrates for a rich Eocene, Antarctic marine environment, in *Paleoenvironment Evolution of Antarctica and the Southern Oceans*, edited by J. Kennett and D. Warnke, pp. 119-130, *AGU Ant. Res. Ser.*, 56.
- Case, J.A. 2006. The late Middle Eocene, terrestrial, vertebrate fauna from Seymour Island: the tails of the Eocene Patagonian size distribution. *in* J.E. Francis, D. Pirrie & J.A. Crame (eds.), "Cretaceous-Tertiary High Latitude Palaeoenvironments, James Ross Basin, Antarctica". Geological Society of London, Special Publication, vol. 258, p. 177-186.
- Deconto, R., and D. Pollard (2003), Rapid Cenozoic glaciation of Antarctica induced by declining atmospheric CO₂, *Nature*, 421, 245-249.
- Dutton, A. L., K. C. Lohmann and W. J. Zinsmeister (2002), Stable isotope and minor element proxies for Eocene climate on Seymour Island, Antarctica, *Paleoceanography*, 17(2), 1-14.
- Francis, J. E., A-M. Tosolini, and D. J. Cantrill (2003), Biodiversity and climatic change in Antarctic Paleogene floras, in *Antarctic Contributions to Global Earth Sciences*, edited by D. K. Fütterer, D. Damaske, G. Kleinschmidt, H. Miller and F. Tessensohn, p. 107, Proceedings of the 9th International Symposium on Antarctic Earth Sciences (ISAES IX), Potsdam, Germany.
- Gandolfo, M. A., P. Hoc, S. N. Santillana, and S. A. Marenni (1998a), Una flor fosil morfológicamente afin a las Grossulariaceae (Orden Rosales) de la Formacion La Meseta (Eocene medio), Isla Marambio, Antartida, in *Paleogeno de America del Sur y de la Peninsula Antartica*, edited by S. Casadio, pp. 147-154, Association Paleontologica Argentina Publicacion Especial, 5.
- Gandolfo, M. A., S. A. Marenni, and S. N. Santillana, (1998b), Flora y paleoclima de la Formacion La Meseta (Eocene medio), Isla Marambio, Antartida. in *Paleogeno de America del Sur y de la Peninsula Antartica*, edited by S. Casadio, pp. 155-162, Association Paleontologica Argentina Publicacion Especial, 5.
- Gradstein, F., J. Ogg, and A. Smith (2004), *A geologic time scale 2004*, Cambridge University Press, Cambridge, UK.
- Ivaney, L. C., S. Van Simaeys, E. W. Domack, and S. D. Sampson (2006), Evidence for an earliest Oligocene ice sheet on the Antarctic Peninsula, *Geology*, 34(5), 377-380.
- Haq, B. U., J. Hardenbol, and P. R. Vail (1987), Chronology of fluctuating sea levels since the Triassic, *Science*, 235, 1156-1167.
- Livermore, R., A. Nankivell, G. Eagles, and P. Morris (2005), Paleogene opening of Drake Passage, *Earth Planetary Sci. Lett.*, 236, 459-470.
- Marenni, S. A. (2006), Eustatically controlled sedimentation recorded by Eocene strata of the James Ross Basin, in *Cretaceous-Tertiary High Latitude Palaeoenvironments*, James Ross Basin, Antarctica, edited by J. E. Francis, D. Pirrie and J. A. Crame, pp. 125-133, *Geol. Soc. Lond. Spec. Publ.*, 258.
- Marenni, S. A., and S. N. Santillana (1994), Unconformity bounded units within the La Meseta Formation, Seymour Island, Antarctica: a preliminary approach, in *Proceedings of the XXI Polar Symposium*, edited by M. Zalewski, pp. 33-37, Institute of Geophysics of Polish Academy of Sciences, Warsaw.
- Poole, I., and D. J. Cantrill (2006), Cretaceous and Cenozoic vegetation of Antarctica Integrating the fossil wood record, in *Cretaceous-Tertiary High Latitude Palaeoenvironments*, James Ross Basin, Antarctica, edited by J. E. Francis, D. Pirrie and J. A. Crame, pp. 63-81, *Geol. Soc. Lond. Spec. Publ.*, 258.
- Reguero, M. A. and S. A. Marenni (in press), Paleogene climatic and biotic events in the terrestrial record of the Antarctic Peninsula: an overview, in *The Paleontology of Gran Barranca: Evolution and Environmental Change through the Middle Cenozoic of Patagonia*, edited by R. H. Madden, A. A. Carlini, M. G. Vucetich, and R. F. Kay, Cambridge University Press, Cambridge, UK.
- Sadler, P. M. (1988), Geometry and stratification of uppermost Cretaceous and Paleogene units on Seymour Island, northern Antarctic Peninsula, in *Geology and Paleontology of Seymour Island, Antarctic Peninsula*, edited by R. M. Feldmann and M. O. Woodburne, pp. 303-320, *Geol. Soc. Am. Mem.*, 169.
- Scher, H. D., and E. E. Martin (2006), Timing and climatic consequence of the opening of Drake Passage, *Science*, 312, 428-430.
- Torres, M. T., S. A. Marenni, and S. N. Santillana, (1994), Maderas fosiles de la Isla Seymour, Formacion La Meseta. Serie Cientifica Instituto Antartico Chileno, 44, 17-38.
- Woodburne, M. O., and J. A. Case, (1996), Dispersal, vicariance, and the Late Cretaceous to Early Tertiary land mammal biogeography from South America to Australia, *Journal of Mammalian Evolution*, 3, 121-161.
- Zachos, J., M. Pagani, L. Sloan, E. Thomas, and K. Billups, (2001), Trends, rhythms and aberrations in global climate 65 Ma to present, *Science*, 292, 686-693.