



Preliminary Model of Porphyry Copper Deposits

Open-File Report 2008–1321

U.S. Department of the Interior
U.S. Geological Survey

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By Byron R. Berger, Robert A. Ayuso, Jeffrey C. Wynn, and Robert R. Seal

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Suggested citation:

Berger, B.R., Ayuso, R.A., Wynn, J.C., and Seal, R.R., 2008, Preliminary model of porphyry copper deposits: U.S. Geological Survey Open-File Report 2008–1321, 55 p.

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Introduction

Porphyry copper deposits result from the complex interactions and feedbacks of many processes. Owing to the complexity, a succinct definition that includes the essential attributes of this deposit class is elusive. Consequently, this descriptive model uses a working definition of a porphyry copper deposit. A porphyry copper deposit is defined as follows:

1. One wherein copper-bearing sulfides are localized in a network of fracture-controlled stockwork veinlets and as disseminated grains in the adjacent altered rock matrix;
2. Alteration and ore mineralization at 1–4-km depth are genetically related to magma reservoirs emplaced into the shallow crust (6–8+ km), predominantly intermediate to silicic in composition, in magmatic arcs above subduction zones;
3. Intrusive rock complexes that are emplaced immediately before porphyry deposit formation and that host the deposits are predominantly in the form of upright-vertical cylindrical stocks and(or) complexes of dikes;
4. Zones of phyllic-argillic and marginal propylitic alteration overlap or surround a potassic alteration assemblage; and,
5. Copper may also be introduced during overprinting phyllic-argillic alteration events.

The general attributes of porphyry copper deposits are summarized below.

Brief Description

In hypogene parts of porphyry copper deposits, the copper occurs predominantly in chalcopyrite; other important copper ore minerals may include bornite and enargite. Ore minerals occur as disseminations and in stockworks of veins in hydrothermally altered, shallow intrusive complexes, often porphyritic, and in adjacent country rocks. This model covers all porphyry-style copper deposits and includes copper-molybdenum, copper-molybdenum-gold, and copper-gold subtypes that are sometimes distinguished by economic geologists (for example, Sillitoe, 2000). A brief discussion of the porphyry copper subtypes is in the Appendix.

Associated Deposit Types

A variety of deposit types are spatially, if not genetically, related to porphyry copper mineralization, including skarns, polymetallic veins and replacements, and epithermal veins. Copper skarn deposits occur near many porphyry copper host intrusions that intruded carbonate-bearing units (Einaudi and others, 1981), and skarn mineral zoning patterns may be useful in the targeting of a potentially associated porphyry copper deposit (Meinert and others, 2005). Typically, the garnet/pyroxene ratio

increases toward the causative skarn-forming pluton with the distal pyroxene zones containing more iron- and manganese-rich pyroxenes than proximal zones (Meinert and others, 2005). In some districts (for example, Christmas, Arizona, and Battle Mountain, Nevada), more copper is recovered from calc-silicate rocks than from the associated intrusive rocks.

Polymetallic replacement deposits occur in carbonate-bearing units peripheral to porphyry-style mineralization. At Bingham, Utah (Babcock and others, 1995), and Bisbee, Arizona (Bryant and Metz, 1966), polymetallic replacement deposits surround the intrusive complexes with offshoots appearing to radiate outward from the stocks. Vein deposits occur peripheral to many porphyry copper deposits (for example, Bingham, Utah), as well as crosscutting porphyry-style mineralization (for example, Valea Morii, Romania).

Primary Commodities

Copper is the primary commodity of economic interest in most porphyry copper deposits, although some deposits with low to very low copper grades are mined principally for their gold (for example, Çöpler, Turkey) and/or molybdenum (for example, Sierrita, Arizona; Continental [Butte], Montana) resources.

By-Product Commodities

Molybdenum, gold, and silver are the economically most important by-products.

Trace Constituents

Rhenium and platinum-group metals are recovered from some deposits. Augé and others (2005) reported the occurrence of platinum-group-metal-bearing tellurides and arsenides in an assemblage of magnetite, chalcopyrite, and bornite in the porphyry deposit at Elatsite, Bulgaria. Tin has been reported from some deposits. Jambor and Owens (1987) found complex tin-bearing sulfide minerals in late-stage veins in the Maggie deposit in British Columbia (Canada), together with a polymetallic assemblage of copper-, zinc-, lead-, and silver-bearing sulfide minerals.

Example Deposits

There are many important deposits, worldwide (compare Singer and others, 2008). Examples of well-studied giant porphyry copper deposits include El Teniente and El Salvador in Chile, Bajo de la Alumbrera in Argentina, Grasberg in Indonesia, and Bingham in the United States.

Regional Environment

Geotectonic Environment

Porphyry copper deposits form in continental magmatic arcs along convergent plate-margin boundaries or in island-arc environments. The deposits are associated with subduction-related volcanic centers, although in some examples, they are thought to be associated with postcollisional volcanism. Permissive magmatic-arc environments may be transpressional or transtensional.

Deformation of magmatic arcs can be exceedingly complex owing to the large variety of mechanisms by which strains can be partitioned in the postmagmatic geologic history of a terrane. Despite the complexity, however, Hindle and Kley (2002) found in the Andes that different scales of analysis elucidate different tectonic effects. For example, near the surface, multitudes of small displacement structures may be typical; yet, at depth, the structures merge and transfer their displacements to large basement faults. Thus, the localization of volcanic centers and their spatial histories may reflect local manifestations of tectonic strains, while chains of volcanic centers may reflect more regional-scale strains. Such issues are important when trying to interpret the localization of porphyry copper deposits within extensive magmatic arcs.

Temporal (Secular) Relations

Owing to the shallow depths of deposit formation (1–4 km), preserved deposits are predominantly Mesozoic and Cenozoic, although there are important older examples. In their worldwide database of porphyry copper deposits, Singer and others (2008) tabulated known or inferred ages for each tabulated deposit.

Relations to Structures

It is generally accepted that porphyry copper deposits are structurally controlled, but opinions vary as to the relative importance of regional compared to local fault systems. Titley (1993) concluded that the importance regional structures have in the localization of deposits is an unresolved issue. Also evaluating the importance of regional-scale structures, Sillitoe (1993) concluded that, at many deposits, there is a lack of evidence that faults of “regional dimensions” control the localization of deposits. More recently, Tosdal and Richards (2001) and Sillitoe and Perelló (2005) also concluded that there are no unique structural environments within which deposits form.

Deposits form in areas of shallow magmatism within subduction-related tectonic environments. The study of a number of deposits (Berger and Drew, 1998; Drew, 2006) shows that there are systematic structural relations that economically viable deposits hold in common. Deposits only form when and where structurally controlled permeability is tightly constrained by regional structure. For example, a common location for deposits is along fault zones where strike-slip displacement is interrupted and transferred onto sets of normal to normal-oblique faults that define an extensional stepover to another, parallel strike-slip fault (also called a pull-apart structure). Fluid flow and magmatism are concentrated into the hinge zones (R. Goldfarb, written commun., 2008), such as at the well-studied Grasberg deposit, Indonesia (for example, Sapiie and Cloos, 2004).

Relations to Igneous Rocks

Porphyry copper deposits result from the condensation of supercritical fluids derived from a crystallizing magma reservoir or set of linked reservoirs in the shallow crust; most likely, the source reservoirs are at depths of 8–10 km or more. A recent study of the Questa porphyry molybdenum deposit in New Mexico suggests that the ore fluid was probably derived from a source reservoir or reservoirs as deep as the middle crust (Klemm and others, 2008). Eight- to 15- or 20-kilometer-deep supercritical ore fluids separate and rise from their sources to form deposits in shallow subvolcanic intrusive complexes at 1–4 km depth. These shallow subvolcanic complexes are typically made up of multiple intrusions of varying composition, as summarized in the “Petrology of Associated Igneous Rocks” section.

For porphyry copper deposits in general, the compositions of the shallow intrusions that host porphyry copper deposits indicate that the deep melt reservoirs from which both the shallow intrusions and the ore fluids were derived may be compositionally calc-alkaline, alkali-calcic, or alkaline (Seedorff and others, 2005). Representative examples of the rock types found in porphyry copper deposits are listed in table 1.

Relations to Sedimentary Rocks

Volcaniclastic rocks are common in many districts, but any sedimentary rock may be spatially associated with a porphyry copper deposit. Where carbonate rocks or carbonate-bearing units are host rocks, replacement skarn assemblages are found.

Relations to Metamorphic Rocks

Metamorphism does not contribute to the formation and localization of porphyry copper deposits, although metamorphosed rocks are host rocks in some districts.

Physical Description of Deposit

Dimensions in Plan View

Singer and others (2008) tabulated the dimensions, including the area of hydrothermal alteration, orebodies, and sulfide-bearing rock, for selected deposits worldwide. These data are shown in table 2. The median size of the longest axis of alteration surrounding a porphyry copper deposit is 4–5 km, while the median size area of alteration is 7–8 km² (table 2).

Vertical Extent

The vertical extent of ore is dependent upon the lower copper cutoff grade, which will vary with the price of copper and local mining costs. At Dos Pobres, Arizona, from the top of the orebody to the 0.2 percent copper cutoff at its deepest points is approximately 1–1.2 km (Langton and Williams, 1982). At Sierrita, Arizona, from ore exposed at the surface to the base of “low

grade” copper mineralization is from approximately 0.8–1.1 km (West and Aiken, 1982). At Bingham, Utah, the greatest vertical extent of copper ore is approximately 1.4 km (Babcock and others, 1995).

Form/Shape

Deposits are typically semicircular to elliptical in plan view. Areal dimensions for ore and alteration (table 2) were calculated assuming an elliptical shape for each deposit unless published information indicated otherwise. Refer to Singer and others (2008) for more information.

In cross section, ore-grade material in a deposit typically has the shape of an inverted cone with the altered, but low-grade, interior of the cone referred to as the “barren” core. In some deposits, for example Grasberg, Indonesia, the low-grade core is a late-stage intrusion (MacDonald and Arnold, 1994).

Host Rocks

Most porphyry copper deposits are associated with shallow intermediate to silicic intrusive complexes composed of small plutons and dikes; some are associated wholly with dikes. Representative host-rock types are listed in table 1. Ore is sometimes hosted in volcanic and sedimentary rocks into which the shallow intrusive rocks were emplaced.

Structural Setting(s) and Controls

There is no broad consensus on the structural localization of porphyry copper deposits at regional scale (for example, Sillitoe, 1993; Titley, 1993). Within deposits, stockwork fracture networks are prominent ore-controlling structures in porphyry-style deposits, and studies such as that by Heidrick and Titley (1982) have shown the fractures to have preferred orientations within a given deposit.

Geophysical Characteristics

Regional

Porphyry copper districts commonly coincide with magnetic anomalies transverse to the predominant structural grain whose gradients represent large regional faults such as near deposits in northern Chile (Behn and others, 2001; Gow and Walshe, 2005), central Iran (Shahabpour, 1999), and New Guinea (Gow and Walshe, 2005).

Local

Porphyry copper deposits often, but not always, appear as magnetic highs, with alteration halos usually manifested as annular (donut-shaped) or open-ring peripheral magnetic lows (Heithersay and Walshe, 1995; Ford and others, 2007). Typically, there is significant variability in magnetic susceptibility throughout the altered rock owing to the nonhomogeneity of phyllic

alteration-related magnetite destruction and late-stage magnetite formation (Gettings, 2005). Porphyry copper deposits almost always appear as moderate gravity lows, especially if the host rock is igneous or metamorphic (Oldenburg and others, 1997). Low, near-surface seismic velocities in porphyry systems correlate well with the phases of a batholith hosting the mineralization, which structurally lie in faulted and brecciated regions (Roy and Clowes, 2000). Mineralized rock almost always has lower resistivity than barren parts of the host stock and surrounding rocks due to the presence of clay minerals and stockwork veins with higher water content, owing to increased fracture permeability.

Induced polarization (IP) anomalies are generally, but not always, a diagnostic indicator of economic mineralization. The IP anomalies correlate with both mineralization and alteration-related magnetic lows; however, IP anomalies often indicate the most abundant pyrite zones in altered rocks rather than areas of less-IP-reactive clay minerals, chalcopyrite, and bornite. Spectral IP has been used to classify different alteration zones and to distinguish noneconomic sulfides such as pyrite from chalcopyrite and bornite (Zonge and Wynn, 1975; Zonge and others, 2005). Resistivity and IP anomaly strengths correlate inversely with resistivity of the host rock and the thickness of any cover. Radiometric methods will show the potassic alteration if significant potassically altered parts of the system are exposed (Sinclair, 1995). Likewise, bands 4/5 and 7/9 of ASTER satellite imagery and radiometric spectrometers have been used to map potassium and thorium anomalies representative of alteration halos surrounding deposits (Rowan and others, 2003; Ranjbar and others, 2004; Shayestehfar and others, 2005).

Potassic, phyllic, argillic, and propylitic alteration halos of porphyry copper deposits contain distinct spectral absorption features that can be mapped using multispectral and hyperspectral remote sensing data (Berger and others, 2003; Mars and Rowan, 2006; Rowan and others, 2006). Argillic minerals alunite, and kaolinite have doublet absorption features at 2.17 and 2.20 μm , and pyrophyllite has a narrow, intense absorption feature at 2.17 μm (Salisbury and others, 1983, and Clark and others, 2007). Phyllically altered rocks typically contain sericite, which is composed of muscovite that has a strong 2.20- μm absorption feature. Propylitically altered rocks contain varying amounts of chlorite, calcite, and epidote, which have 2.32–2.35- μm absorption features (Salisbury and others, 1983; Rowan and others, 2006; Clark and others, 2007). In the thermal infrared region, quartz and hydrous silica contain emission minima near 8.40 and 8.95 μm , which are typically found in potassic-altered cores that have been enriched in silica (Rowan and others, 2006).

Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) multispectral data have been used to regionally map porphyry copper deposits in Iran, Afghanistan, and Kazakhstan (Mars and Rowan, 2006). Argillically altered rocks (containing alunite and kaolinite) of porphyry copper systems were regionally mapped using ASTER band ratios b4/b5, b5/b6, and b6/b7 combined in a logical operator algorithm (Mars and Rowan, 2006). Using the same band ratios and methods, but different band thresholds, muscovite-rich, phyllic-altered rocks have also been regionally mapped (Mars and Rowan, 2006). In regional studies, porphyry copper deposits that are not structurally deformed are typically associated with semicircular to elliptical patterns of argillically and phyllically altered rocks (Mars and Rowan, 2006). Siliceous, phyllic, argillic, and propylitic rocks have been mapped using ASTER data and a matched filtering algorithm incorporating image spectra at the Reko Diq porphyry copper deposit in Pakistan (Rowan and others, 2006). Siliceous rocks were also mapped at Reko Diq, Pakistan, using ASTER band ratio 14/12 (Rowan and others, 2006).

The hyperspectral Airborne Visible Infrared Imaging Spectrometer (AVIRIS) has also been used to map altered rocks associated with porphyry copper deposits at the Red Mountain and Sunnyside deposits in the Northern Patagonia Mountains of

Arizona (Green and others, 1998; Berger and others, 2003). Alunite, kaolinite, pyrophyllite, and muscovite were mapped at the Red Mountain and Sunnyside deposits using Tetracorder, a spectral feature-fitting algorithm (Clark and others, 2003).

Ore Characteristics

Mineralogy

The ore mineralogy of selected porphyry copper deposits is given in table 3. The principal hypogene, copper-bearing sulfide mineral is chalcopyrite, although substantial amounts of copper may occur as bornite, enargite, and chalcocite. By-product minerals frequently include molybdenite and native gold. Other associated minerals may include sphalerite, galena, tetrahedrite, and gold tellurides.

Mineral Assemblages

Copper-ore mineral assemblages are a function of the chemical composition of the fluid phase and the pressure and temperature conditions affecting the fluid. Thus, specific mineral associations may vary in a deposit as a function of space and time as the composition of the hydrothermal fluid changes. In primary, unoxidized or not supergene-enriched ores, the most common ore-sulfide assemblage is chalcopyrite \pm bornite, with pyrite and minor amounts of molybdenite. In some deposits, there is an advanced argillic alteration that overprints near-neutral pH alteration and mineralization. For example, at Rosia Poieni, Romania, the advanced argillic ore-sulfide assemblage is pyrite + enargite + marcasite + chalcocite (Milu and others, 2004). In supergene enriched ores, a typical assemblage might be chalcocite + covellite \pm bornite (Schwartz, 1966). In oxide ores, a typical assemblage might include malachite + azurite + cuprite + chrysocolla, with minor amounts of other minerals (for example, carbonates, sulfates, phosphates, and silicates) (Schwartz, 1966).

Paragenesis

For deposits that have been studied in detail, sulfide-ore mineral paragenesis varies widely within and between deposits. Bagdad, Arizona, illustrates the complexity that is encountered in a porphyry-style deposit, although the specifics of the Bagdad paragenesis should not be construed as a model for all porphyry copper deposits. Barra and others (2003) studied the sulfide mineralogy at Bagdad. They defined “early mineralization” wherein molybdenite was the sole ore sulfide mineral present. Subsequent “hypogene” mineralization paragenesis included assemblages of molybdenite, pyrite + chalcopyrite, molybdenite, pyrite, chalcopyrite + pyrite, molybdenite + pyrite + chalcopyrite, and sphalerite + galena + tetrahedrite \pm silver. This paragenetic sequence differs from that published by Anderson and others (1955) who did not observe early-stage molybdenite mineralization but did observe the later molybdenite + pyrite-bearing veins crosscutting chalcopyrite + pyrite veins. In many cases, the paragenetic sequence reported may be reflective to the date of a study, considering the extent of mining at that date. At Batu Hijau, Indonesia, Arif and Baker (2004) looked in detail at the copper sulfide paragenesis, particularly with respect to its association with gold. The early-stage veins contain bornite, digenite, and chalcocite wherein digenite occurs as exsolution lamellae

in the bornite and chalcocite rims the bornite-digenite grains. Crosscutting veins contain chalcopyrite \pm bornite. Arif and Baker (2004) also note evidence that chalcopyrite \pm pyrite replaced some early-stage bornite + digenite \pm chalcocite. Cubanite has been reported from some porphyry copper deposits (table 3). Ramdohr (1980) suggests that such occurrences, for example the digenite at Batu Hijau, are due to exsolution from an originally higher temperature solid solution.

Zoning Patterns

The zoning of sulfide minerals, particularly with respect to pyrite content, has been documented in many deposits. Nielsen (1968) found that the pyrite content at Santa Rita, New Mexico, increases from irregular zones with <1 wt (weight) percent pyrite in the central part of the deposit, through a surrounding zone with 1–4 wt percent pyrite to an outer zone with >4 wt percent. The greatest chalcopyrite content, >0.4 wt percent chalcopyrite, tended to be in the zone with intermediate amounts of pyrite. Inward from this higher chalcopyrite zone, the chalcopyrite decreased to <0.1 wt percent, whereas outward from the higher grade zone it ranged between 0.1 wt percent and 0.4 wt percent. At Bingham, Utah, John (1978) defined five overlapping, but concentric, zones as follows from the center outward: deep low-grade core (<0.5 wt percent sulfides including pyrite, chalcopyrite, bornite, and molybdenite), molybdenite, bornite + chalcopyrite, chalcopyrite + pyrite, pyrite, and galena + sphalerite. As at Santa Rita, at Bingham the inner zone has little pyrite, the copper ores contain 1–2 wt percent pyrite, and surrounding the copper ore the pyrite content can exceed 4 wt percent.

Textures and Structures

Ore minerals occur disseminated through the altered rock matrix and in discrete veins filling fractures. At Bagdad, Anderson and others (1955) observed that some disseminated sulfides form linear chains of minerals that are suggestive that even disseminated mineralization was localized by fractures, now annealed. In the disseminated ores, individual sulfide grains may exceed 1 mm in longest dimension, but most are smaller (Anderson and others, 1955). Ore minerals in discrete veins may occur as dispersed grains or in clots and occur also in alteration selvages to the veins.

Grain Size

Typically, the principal copper sulfides and molybdenite consist of millimeter-scale grains, but may be as large as 1–2 cm in diameter and, rarely, pegmatitic (larger than 2 cm).

Gangue Characteristics

Mineralogy

In porphyry copper deposits, copper ores include disseminated and vein-form styles of mineralization. Thus, gangue material reflects many or all of the sulfide minerals in the “Ore Characteristics” section and the “Alteration” section. The information

in these sections is summarized in table 2. The most abundant gangue sulfide is pyrite. Other gangue sulfides found, but generally in small quantities, include pyrrhotite, sphalerite, tennantite, and galena. The most abundant gangue silicate and oxide minerals are quartz, K-feldspar, biotite, muscovite (sericite), magnetite, chlorite, anhydrite, and epidote. Other commonly occurring gangue minerals are rutile, apatite, and titanite.

Mineral Assemblages

The mineral assemblages are the same as in the “Ore Characteristics” section and the “Alteration” section. As noted by Schwartz (1947) from the study of 12 deposits in North America (table 4), the hydrothermal alteration and mineralization in porphyry copper deposits is complex and variable both within individual deposits and between different deposits.

Paragenesis

The specifics of gangue paragenesis vary within and between deposits. Some examples are given in the “Ore Characteristics” section and the “Alteration” section.

Zoning Patterns

Zoning patterns follow mineralization and alteration zoning patterns as discussed in the “Ore Characteristics” section and the “Alteration” section. Titley (1982) reports that different alteration mineral assemblages develop sequentially and the vein-gangue assemblage zoning reflects this paragenesis.

Textures and Structures

Pervasive alteration results in the partial to complete replacement of the primary host-rock minerals. Phenocrysts in porphyritic rocks are typically pseudomorphed, with the individual alteration grains replacing the phenocrysts consisting of an agglomeration of small crystals that take on the form of the phenocryst. Ragged-appearing biotite commonly occurs in this manner. Linear arrays of gangue minerals, such as pyrite or potassium feldspar, may be present with no apparent fracture control. Only wispy veinlets are generally found in pervasively altered rock.

Stockworks of veins, predominantly quartz, are a common form of gangue. Veins are often classified according to crosscutting relations in conjunction with form (for example, A, B, and D veins of Gustafson and Hunt, 1975). For example, early stage veins in many porphyry copper deposits have irregular or wavy margins, typically lack banding, and have a granular texture. Later stage, crosscutting veins typically have sharp boundaries, and banding is more common.

Grain Size

Grain sizes vary within and between deposits; variation is from microscopic to millimeter to centimeter scale.

Geochemical Characteristics

Trace Elements and Element Associations

Deposit-scale studies of trace elements have shown that copper, molybdenum, gold, silver, barium, lead, zinc, arsenic, antimony, and tellurium make up suites related to alteration zoning within deposits (for example, Beus and Grigorian, 1977; Chaffee, 1982, 1992, 1994). The principal associations are copper-molybdenum-gold \pm silver-arsenic and barium-zinc-lead-antimony-silver-arsenic-tellurium \pm bismuth.

Olade and Fletcher (1976) found that sulfide iron and copper concentrations, as determined from selective leaches, are highest in ore-bearing rock. Total sulfur anomalies broadly outline mineralization and are often highest in pyritic propylitic alteration zones. Similar results were obtained by Chaffee (1982).

Gold is an important by-product in many porphyry copper deposits, but the details of its occurrence are not broadly understood (for example, Arif and Baker, 2004). A comprehensive study of gold was conducted at Batu Hijau, Indonesia, by Arif and Baker (2004). They found gold in quartz veins, attached to sulfide grains, and as free gold along silicate grain boundaries. Gold is predominantly found in association with bornite; and, in bornite-rich ores, the gold is within the sulfide mineral structure and present as small grains of native gold. In chalcopyrite-rich ores, the gold occurs primarily as small grains within the copper sulfide.

Zoning Patterns

A summary of zoning studies was published by Govett (1983). Copper, molybdenum, gold, and tin are most anomalous within the ore zones. Although silver and arsenic are also anomalous within ore zones, they also occur with barium, zinc, lead, and antimony in immediately overlying subeconomic, altered rock (Beus and Grigorian, 1977). Govett's (1983) summary shows that silver, zinc, lead, antimony, arsenic, tellurium, and bismuth are anomalous lateral to ore zones. At Kalamazoo, Arizona, Chaffee (1982) found that tellurium was most anomalous above ore and that zinc and lead form negative anomalies immediately above and within ore.

Fluid-Inclusion Thermometry and Geochemistry

A wide range of fluid-inclusion homogenization temperatures has been obtained from porphyry copper deposits. Klemm and others (2007) reported homogenization temperatures from about 550°C to about 250°C, with single measurements within individual assemblages varying by as much as about 100°C for samples from El Teniente, Chile. When binned in 10°C intervals, the median reported homogenization temperature range in the El Teniente data is 380°–390°C. At Bajo de la Alumbrera, Argentina, Ulrich and others (2001) reported homogenization temperatures varying from about 770°C to about 200°C, with single measurements within individual assemblages varying by as much as 100°C or more. When binned in 10°C intervals, the median reported homogenization temperature range (excluding extrapolated temperatures) in the Bajo de la Alumbrera data is 470°–480°C.

Bodnar (1995) summarized evidence that the source of metals in porphyry deposits is magmatic and described fluid-inclusion evidence. Bodnar and Cline (1991) and Bodnar (1995) described magmatic fluid inclusions as high-salinity and halite-bearing, which homogenize at relatively high temperature (greater than 500°–600°C). Magmatic fluid inclusions are found in quartz in early veins and show a range of characteristics that reflect the different depths at which the inclusions were trapped. For example, the deeper levels of a porphyry system are represented at the Butte, Montana, porphyry copper-molybdenum deposit (Montana) (Roberts, 1975; Rusk and others, 2008), where the chalcopyrite-bearing inclusions in deep veins have moderate to low salinity (2–5 wt percent sodium chloride equiv), 2–8 mol-percent carbon dioxide, and have moderate homogenization temperatures (575°–650°C) and pressures equivalent to about 6–9-km depth (Rusk and others, 2008). In shallower systems, as in the Red Mountain, Arizona, porphyry copper deposit, the magmatic fluids formed vapor-rich inclusions that trapped a low-density, low-salinity fluid that exsolved directly from the magma (Bodnar, 1995). Hedenquist and Richards (1998, and references therein) explained the coexistence of hypersaline-liquid-rich and vapor-rich fluid inclusions by the unmixing of an originally homogeneous aqueous fluid that exsolved from magma at pressures of 1–1.5 kbar. The magmatic fluid is transported to relatively shallow crustal levels, and ore deposition takes place at ambient pressures of about 0.5 kbar where the fluid intersects its solvus, producing immiscible saline liquid and vapor phases.

In the magma reservoir from which the hydrothermal fluids are exsolved, decompression and fractional crystallization of silicic magma triggers separation of an aqueous phase and bubbles in the residual melt (Burnham, 1979). The function of fluid-phase separation in ore-metal fractionation and mineral precipitation resulting from the wide density variations and degree of miscibility of saline fluids between surface and magmatic conditions have been the focus of studies by Heinrich (2007). Evidence for volatile separation includes the presence of miarolitic cavities (Candela and Blevin, 1995) and pods of saccharoidal quartz connected by anastomosing zones of graphic quartz-alkali feldspar intergrowths and ragged biotite (with lesser apatite and magnetite) (Harris and others, 2004). In the Western United States, crystals of copper and molybdenum sulfides have been found in the miarolitic cavities (for example, Wilson, 1975). Gustafson and Hunt (1975) recognized several intrusive stages related to the El Salvador, Chile, porphyry copper deposit, described fluids associated with different generations of veins, and attributed the early-alteration-mineralization stage to magmatic fluids. Dilles (1987) studied the evolution of ore fluids in the Yerington, Nevada, porphyry copper district and concluded that an aqueous fluid, rich in sodium, chloride, potassium, iron, and sulfur species, causes copper and molybdenum sulfide mineralization. Harris and others (2003, 2004) documented the presence of interconnected miarolitic cavities (and associated quartz segregations) and investigated silicate-melt inclusions and coexisting aqueous phases captured in primary inclusions at the Bajo de la Alumbrera, Argentina, porphyry copper-gold deposit. Such textural and silicate-melt inclusion data document volatile-phase separation and link magmatic processes to hydrothermal alteration.

The compositional evolution of the fluids had been constrained by documentation of petrography, microthermometry, and LA-ICP-MS analysis of fluid inclusions (for example, Ulrich and others, 2001; Klemm and others, 2007). Scanning electron microscope cathodoluminescence studies have shown the successive generations of fluid inclusions in texturally complex quartz veinlets during the main stages of metal transfer in the porphyry copper-gold-molybdenum deposit at Bingham (Landtwing and others, 2005). At Bingham, early quartz veins are brightly luminescent and crystallized before the copper-iron sulfides precipitated in these veins. The LA-ICP-MS analyses also show that fluids trapped before and after precipitation of the copper-iron sulfides are largely similar in their contents of major, minor, and trace elements, except for copper (Landtwing and others, 2005).

Alteration

Mineralogy

In porphyry copper deposits, hypogene hydrothermal alteration is typically classified on the basis of mineral assemblages, discussed in more detail herein. In silicate-rich rocks, the most common alteration minerals are K-feldspar, biotite, muscovite (sericite), albite, anhydrite, chlorite, calcite, epidote, and kaolinite. In silicate-rich rocks that have been altered to advanced argillic assemblages, the most common minerals are quartz, alunite, pyrophyllite, dickite, diaspore, and zunyite. In carbonate rocks, the most common minerals are garnet, pyroxene, epidote, quartz, actinolite, chlorite, biotite, calcite, dolomite, K-feldspar, and wollastonite. Other alteration minerals commonly found in porphyry copper deposits are tourmaline, andalusite, and actinolite. Table 3 summarizes alteration minerals in selected porphyry copper deposits, worldwide.

Alteration Mineral Assemblages

Assemblages of alteration minerals are typically grouped in categories (Schwartz, 1947; Creasey, 1966; Meyer and Hemley, 1967; Rose, 1970; Beane, 1982). Potassic alteration generally includes one or more minerals in the assemblage K-feldspar-biotite-muscovite (sericite)-chlorite-quartz. Phyllic alteration includes one or more minerals in the assemblage quartz-muscovite (sericite)-pyrite-chlorite. Argillic alteration includes one or more minerals in the assemblage kaolinite (\pm dickite)-muscovite (sericite)-montmorillonite. The “advanced-argillic” mineral assemblage is quartz-alunite-kaolinite (\pm dickite) \pm pyrophyllite. Propylitic alteration consists of one or more minerals of the assemblage chlorite-calcite-epidote-pyrite-albite.

Hydrothermal alteration reflects changes in the host-rock mineralogy in response to water/rock interactions. Thus, the general categories of alteration (for example, potassic) can consist of different combinations of the minerals previously listed; that is, different assemblages depending upon fluid and host-rock compositions. Because there may be multiple, overlapping pulses of hydrothermal fluid flow that collectively form a deposit, the host-rock composition that each sequential fluid pulse contacts may be a previously altered rock. Thus, assemblages within any particular group may be replicated at different times across a deposit and(or) may differ somewhat in relative constituent mineral abundances from preceding assemblages of the same overall mineralogy. In addition, the assemblage may vary depending upon whether the alteration is pervasive or restricted to the selvage of a vein. The mineral-assemblage complexity that may exist in a deposit is illustrated in table 4.

Lateral and Vertical Dimensions

Table 2 gives the dimensions of the alteration systems for a selection of deposits (from Singer and others, 2008). Lowell and Guilbert (1970) estimated that alteration extends, on average, 0.75 km beyond the limits of ore in deposits in the western Cordillera of the United States.

Vein Selvages

As used herein, selvages refer to distinct bands of wall-rock alteration adjoining veins. Because mineral reactions vary depending upon host-rock composition, vein-margin assemblages can vary widely. Examples of vein selvage alteration for two deposits are given in table 4.

Matrix

Irrespective of the alteration category (for example, potassic), the degree of alteration of the original rock matrix can vary from 100 percent to some smaller proportion, depending upon the stability of the host-rock-forming minerals under the hydrothermal conditions and the intensity of the alteration. Except for extremely intense argillic, advanced argillic, phyllic alteration, the original texture of the host rocks is preserved.

In the potassic alteration zones in quartz-bearing host rocks, the primary quartz generally remains unaltered. In mafic mineral-bearing rocks (for example, biotite, amphibole, pyroxene), secondary biotite may be abundant together with chlorite, sulfides, sericite, a titanium-bearing phase (for example, rutile, ilmenite, sphene), epidote, and(or) calcite. Primary plagioclase is generally altered, particularly the more calcic varieties. The most common alteration phases are K-feldspar, albite-oligoclase, and sericite (compare Beane, 1982). Anhydrite is common, but its abundance varies greatly between deposits.

Intensity

The intensity of alteration can vary from 100 percent of specific primary minerals in the host rocks to partial alteration of susceptible minerals resulting in a somewhat smaller proportion of alteration of the original rock. In the most intensely altered rocks, the primary textures of the host rock may have been largely to entirely destroyed and replaced by mats of fine-grained alteration minerals, particularly biotite and(or) sericite, or, in advanced argillic alteration, mats of alunite and pyrophyllite.

Textures

In pervasively altered rock, alteration minerals tend to be finer grained than their primary antecedents. Phenocrysts are commonly pseudomorphed, but replacement minerals, whether biotite, sericite, chlorite, or some other phase, form aggregates of small grains that mimic the original crystal. Alteration minerals are subhedral to anhedral, except when they form in vugs or veins. Whether the alteration is pervasive or vein related, it is not unusual for pyrite to be euhedral.

Zoning Patterns

Despite the intrinsic complexity of hydrothermal alteration and the significant variability within and between deposits, the same general zoning of the alteration categories is macroscopically evident in most deposits. The zones are defined on the basis of their characteristic minerals as listed in the “Mineralogy” section. The compositional changes observed are the consequence of the involvement of differently sourced fluids and their mixture in the formation of a deposit. Figure 1 is a conceptual model of

the lateral and vertical relations of the alteration zones first proposed by Lowell (1968) and later applied to many deposits in the North American Cordillera (Lowell and Guilbert, 1970).

When a deposit begins to form, there are two distinct zones that form. In the core of the upflow zone of the mineralizing magmatic fluid, potassic alteration occurs. With respect to the deposition of copper ores, potassic alteration in the core of a deposit is of predominant importance. Stable isotopic evidence, discussed in a later section, indicates that these ore-forming, magmatic fluids do not mix with surrounding ground waters. Nevertheless, the thermal gradient associated with this high-temperature upflow zone leads to convection of surrounding ground waters that results in a peripheral propylitic alteration zone. Phyllic alteration is always observed to crosscut potassic alteration, and isotopic evidence, discussed in a later section, indicates that this alteration forms from a mixture of meteoric and magmatic fluids. Phyllic alteration is associated with important ton-nages of ore in some deposits but is not present as a distinct alteration type in all deposits (Sillitoe, 2000).

Clay-rich alteration assemblages, such as argillic to advanced argillic, commonly occur above the core of a deposit and laterally along the margins of the system. This upper alteration zone is sometimes referred to as a “lithocap” (for example, Sillitoe, 2000).

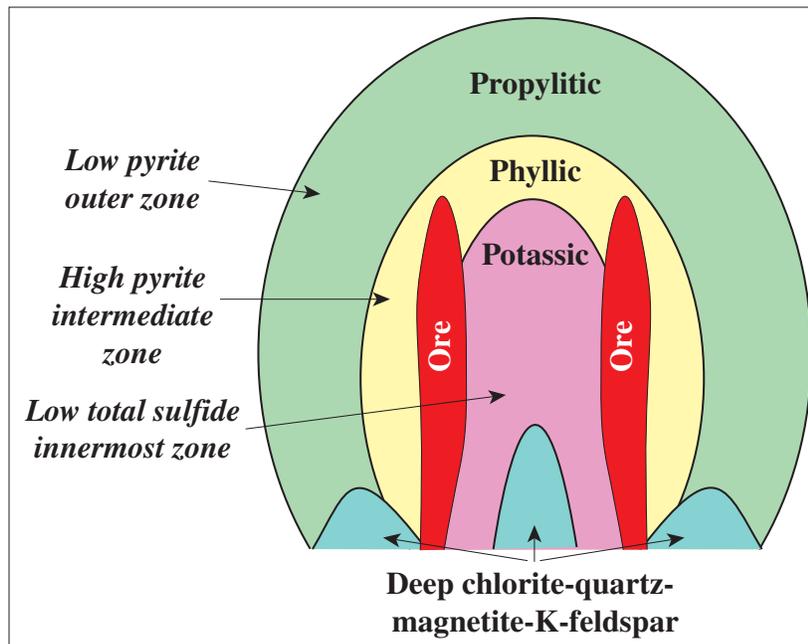


Figure 1. Cross section of a porphyry copper deposit showing idealized alteration zoning (after Lowell and Guilbert, 1970).

Petrology of Associated Igneous Rocks

Rock Names

A wide variety of igneous rock types, depending on the depth and chemical conditions of magma generation, are spatially associated with porphyry deposits, ranging from diorite to granodiorite and granite. The most common rock association

is related to the evolution of felsic to intermediate calc-alkaline compositions (granodioritic), likely derived from more mafic mantle magmas. In island-arc settings, the crust is thinner and the host rocks are more mafic (dioritic), and in back-arc settings the host rocks are more alkaline (shoshonitic) (Richards, 2003).

Forms of Igneous Rocks and Rock Associations

Most immediately associated intrusions are small stocks, often with a greater vertical than horizontal dimension, and dikes. Other igneous rocks can vary from dikes to small stocks to batholith-scale bodies.

Mineralogy

Most common primary minerals in host quartz dioritic to granodioritic rocks include quartz, K-feldspar, plagioclase, biotite, hornblende, magnetite, zircon, apatite, and titanite. Porphyry copper-gold deposits are genetically associated with the magnetite series of granites, as defined by Ishihara (1981, and references therein). The oxidation state of the magma is largely inherited from the source region (Carmichael, 1991). Other parameters, such as magma compositional variations (for example, its degree of peraluminosity, range of alkalinity, silica, water content, and so forth) are also inherited from the source region. The compositional features and degree of magma evolution control the compatible/incompatible element distributions (Blevin and Chappell, 1992).

Textures and Structures

Most immediately associated igneous rocks are porphyritic to microporphyritic with aplitic groundmass; dikes may have aphanitic groundmass. In some deposits, minerals show preferred orientations related to flow or postemplacement deformation.

Grain Size

Grain sizes can vary from coarse to aphanitic but are predominantly intermediate to fine grained.

Lithochemistry

Porphyry copper-(molybdenum-gold) deposits are formed by magmatic-hydrothermal fluids generated from subduction-related magmatism. Calc-alkaline igneous rocks associated with porphyry copper deposits vary from predominantly hornblende- and/or biotite-bearing diorite to monzogranite (Cooke and others, 1998, 2005; Richards, 2003; Sinclair, 2007; Sillitoe, 1998; Sillitoe and Perelló, 2005, and references therein; Seedorff and others, 2005, and references therein). Syenogranite, quartz monzonite, and quartz monzodiorite occur in some porphyry copper deposits, and diorite and pyroxenite have also been reported from more mafic and alkalic varieties (for example, Cooke and others, 1998; Panteleyev, 1995a, b). High-potassium calc-alkalic (and alkalic) intrusions are related to gold-rich porphyry systems (for example, Cooke and others, 2005). The more mafic end of the granitic compositional spectrum is closely related to copper (-gold) mineralization (for example, Blevin and Chappell, 1992, 1995). Deposits are associated with multiphase, shallow (porphyritic rocks with aplitic groundmass), moderately evolved

granitic rocks (as judged by chemical parameters such as K/Rb, and Rb/Sr ratios, and moderate silica contents), sulfur-rich, and oxidized (high values of $\text{Fe}_2\text{O}_3/\text{FeO}$ in bulk rock; belonging to magnetite-series, titanite+magnetite-bearing granitic rocks) magmatic systems (intrusions and volcanic rocks). Seedorff and others (2005) found that porphyry deposits in the Western United States are spatially, temporally, and genetically related to metaluminous to weakly peraluminous, and intermediate to silicic stocks ($\text{SiO}_2 > 56$ wt percent). White (2004) suggested that porphyry copper deposits in the Western United States were associated predominantly with high-temperature quartz monzonite characterized by relatively high contents of $\text{K}_2\text{O} + \text{Na}_2\text{O}$, $\text{K}_2\text{O}/\text{Na}_2\text{O} > 1$, and $\text{Ba} > 1,000$ ppm.

Trace-Element Geochemistry

The modern level of understanding of the chemical evolution of magmatic systems and porphyry copper deposits is not precise enough to predict the exact geochemical features that distinguish mineralized from barren plutons. However, important advances have been made in some cases. For example, Lang and Tittley (1998) found that rare-earth element (REE) chondrite-normalized patterns of productive plutons in Arizona have steeper slopes, lower total REE contents, increasing upward concavity in the heavy-REE, and less negative to positive europium anomalies relative to nonproductive stocks; the high-field-strength elements (zirconium, hafnium, tantalum, niobium), and manganese and yttrium are depleted in productive stocks compared to nonproductive stocks. Baldwin and Pearce (1982), Kay and others (1999), Richards and others (2001), and Richards (2003) noted compositional distinctions in the Andes, including productive intrusions associated with porphyry deposits that show fractionated REE patterns and depletions for manganese, thorium, yttrium, and the heavy-REE compared to barren plutons. Major porphyry deposits in the Peruvian-Chilean belt are also known to have formed late within a magmatic cycle, hosted by more chemically evolved rocks that are assigned to the last intrusive event in the region (for example, host dacites evolving from cogenetic diorites; Richards, 2003 and references therein). In such cases, the heavy-REE and yttrium depletions are attributed to amphibole and(or) deep-crustal garnet fractionation and the low manganese to loss of manganese-rich fluids from the magma.

Stable Isotope Geochemistry

Early work by Sheppard and colleagues (for example, Sheppard and others, 1969; Sheppard and Gustafson, 1976) using oxygen and hydrogen isotopes showed that magmatic fluid dominates the early, high-temperature stages of porphyry evolution and formation of biotite \pm K-feldspar (potassic alteration); a component of meteoric water subsequently overprints the early hydrothermal stages producing muscovite \pm pyrite assemblages (phyllic alteration) and clay-bearing assemblages (argillic alteration) at lower temperatures (Hedenquist and Richards, 1998; Seedorff and others, 2005, and references therein). Magmatic fluids dominate acidic alteration associated with ore (for example, Watanabe and Hedenquist, 2001). Ohmoto and Rye (1979) found that sulfides and sulfates in porphyry copper deposits from the American Cordillera were deposited from fluids carrying sulfur of predominantly magmatic origin.

Radiogenic Isotope Geochemistry

The typical association of porphyry copper deposits and ordinary calc-alkaline magmas points to commonly accepted and standard genetic processes of magma generation in magmatic arcs involving melting of the metasomatized mantle wedge (Richards, 2003). In many study areas, the current level of knowledge of the tectonic history, age relationships, and structural evolution are not sufficient to recognize regionally systematic isotopic signatures that can be attributed to diagnostic mantle or crustal sources forming porphyry copper deposits. Moreover, calc-alkaline magmas that have been thought to be intrinsically capable of producing copper-rich fluids and porphyry copper deposits appear to require no exceptional sources (Dilles, 1987; Cline and Bodnar, 1991).

Recent studies have shown, however, large-scale regional isotopic variations ascribed to source or to magmatic processes during transfer through the crust. For example, studies of lead-strontium-neodymium-osmium isotopes link deposits in the Andes to mantle-derived magmas with contributions from the continental crust (Hedenquist and Richards, 1998; Ruiz and Mathur, 1999; Sillitoe and Perelló, 2005); in southwestern Arizona, the isotopes reflect a large contribution from the continental crust (Anthony and Titley, 1988; Lang and Titley, 1998). Homogeneous lead isotope ratios characterize individual eastward-younging metallogenic belts in the Andes (Tosdal and others, 1999, and references therein), in contrast to regionally heterogeneous ratios in southwestern Arizona (Wooden and others, 1988; Bouse and others, 1999). Neodymium and strontium isotope ratios do not distinguish between mineralized and barren systems in North America (Farmer and DePaolo, 1984). The osmium isotope ratios point to the relative contributions of copper from the mantle and crust (Ruiz and Mathur, 1999). In the larger deposits in Chile, the initial osmium isotope ratios are less radiogenic, suggesting that such deposits acquire relatively more osmium from the mantle (Mathur and others, 2000). Long-term underplating of basalt and recycling of mafic crust produced significantly thickened crust in the Andes; systematic changes in neodymium-strontium-lead ratios from west to east have also been found (Haschke and others, 2002).

Depth of Emplacement

We believe that most deposits formed between about 1.5- and 3–4-km depth. Other investigators have estimated some formation depths in the range between 4–5 (Sutherland Brown, 1976) and 6–9 km (Rusk and others, 2008). A summary of published estimates is given in Singer and others (2008).

Theory of Deposit Formation

System Affiliation(s)

Seedorff and others (2005) noted that porphyry copper deposits can be associated with high-sulfidation epithermal deposits (copper, gold, silver), late and/or distal intermediate-sulfidation polymetallic base metal and precious element veins (lead, manganese, zinc, silver), and distal disseminated gold deposits. Sillitoe and Perelló (2005) described iron-oxide-copper-gold, manto copper, copper-bearing volcanogenic massive sulfide deposits, and copper skarns (copper-zinc-molybdenum-silver) spatially associated with porphyry copper belts in the Andes. Jensen and Barton (2000) described alkalic porphyry systems with

intermediate-sulfidation epithermal base-metal-gold-telluride deposits. Porphyry copper deposits can exhibit regional or district-scale zoning, from a porphyry core associated with proximal skarns to distal polymetallic veins and replacement deposits.

Controls on Permeability and Fluid Flow

Deposits may be broadly related to regional and pull-apart structures at dilational bends, strike-slip faults, shear zones, duplexes, pull-apart basins, and grabens. At the local scale, extensive sets of fractures develop in response to hydraulic fracturing.

Sources of Fluids and Metals

The calc-alkaline nature of the magmas reflects melting of various sources (mafic rocks) including those in the mantle wedge (modified by metasomatism of fluid-soluble elements and melts during subduction), magma ponding and reaction at the base of the continental crust (hybridization, mafic and felsic magma mixing), and metasomatism during ascent through thickened crust (for example, Richards, 2003). These fundamental aspects of magma production in continental margins point to multiple, multistage processes that were first summarized by Hildreth and Moorbath (1988) as the melting, assimilation, storage, and homogenization (MASH) model to explain crustal contributions to Andean arc magmatism. Such a diversity of magma sources and multistage igneous processes invoked for magma genesis suggests that multiple metal sources are involved in porphyry copper deposits. The mantle likely provides most of the copper, gold, and platinum-group elements, and the crust provides most of the molybdenum and lead.

Chemical Transport and Transfer Processes

Fluid-phase separation in porphyry copper deposits controls ore-metal fractionation and sulfide mineral precipitation and results from density variations and degree of miscibility of saline fluids (Heinrich, 2007). Extensive magma evolution in the shallow crustal environment produces supersaturation of volatile phases and an aqueous phase that ponds in the periphery and above the crystallizing magma chamber. Aqueous fluid has commonly been thought to be the main agent of metal transport in porphyry copper deposits, but the function of vapor transfer was first proposed for the formation of porphyry deposits by Henley and McNabb (1978). Williams-Jones and Heinrich (2005) recently evaluated the function of aqueous vapor as an agent of metal transport in natural systems and in the laboratory and concluded that vapor, instead of aqueous fluid, is the main transporting agent. Other recent studies of porphyry copper deposits have also shown that copper and gold partition into the vapor (Heinrich and others, 1999; Ulrich and others, 1999). Parental magmas producing porphyry copper deposits may not be highly crystallized or represent an advanced stage of evolution. Sinclair (2007) summarized evidence for highly reactive volatile streams originating deep within a magma reservoir (more mafic zones? separate reservoirs?), acquiring metals during transport, and ponding in cupolas (without triggering an eruption). These large volumes of volatiles and ore-forming fluids strip metals during ascent, depressing the liquidus temperatures of granitic magmas in the cupolas. Compositionally zoned magma reservoirs and caldera processes involving resurgence of less felsic, deeper, hotter magmas appear to be involved with the generation of highly reactive, volatile streams. Richards (2003) suggested that caldera and ignimbrite complexes are probably not prospective for porphyry

copper deposits because large caldera-forming eruptions are expected to obliterate magmatic-hydrothermal systems, and ignimbrites predominantly reflect crustal melts that are relatively depleted in sulfur and chalcophile metals.

Heat Transport and Transfer Processes

Cooling of magmatic-hydrothermal fluids from about 750° to 600°C (depending on magmatic composition) to less than 200°C (Ulrich and others, 2001) promotes supersaturation of ore and alteration minerals (Hemley and Hunt, 1992).

Pressure Gradients during Ore Formation

Magma reservoirs are dynamic and chemically open systems that de-gas, assimilate crust, recharge with new fluxes of mafic melts at their base, and mix different compositional zones during magma evolution. Magma composition and water content, together with the depth of the reservoir and hydrologic regime, establish important constraints on ore deposition; near the magma reservoir at high temperature, overpressured and lithostatic conditions occur, but at temperatures less than 375°C, hydrostatic conditions dominate (Seedorff and others, 2005). Multiple batches of porphyry magma (dikes, plugs) intrude into the ore zone, possibly creating breccias and hydraulic fractures. At low pressures (<1.5 kbars) brine-vapor-phase separation controls metal partitioning and solubility (Heinrich and others, 1999; Heinrich, 2007).

Wall-Rock Reaction and Ore-Mineral Deposition Processes

Wall-rock composition affects the alteration and ore mineral assemblage by buffering the pH and oxidation state of the system.

Weathering

Weathering, resulting in oxidation and(or) supergene enrichment, is an important process in the economic viability of many porphyry copper deposits (for example, Sillitoe and McKee, 1996). The oxidation and acid leaching of primary mineralization may produce zones of enrichment near the base of the weathering zone (Hartley and Rice, 2005). Primary rock composition and postmineralization hydrogeology, together with one or more cycles of uplift and(or) climate change, are important to the formation of an economically viable deposit. Prominent examples of oxidation and supergene enrichment include the deposits at Morenci, Arizona, and Chuquicamata, Chile.

The mineralogy of weathered deposits has been discussed in detail by Anderson (1982), Alpers and Brimhall (1989), and Sillitoe (2005). Common minerals in strongly leached rocks, referred to as “leached caps,” include goethite, jarosite, hematite, gypsum, alunite, and kaolinite. Supergene and oxide minerals include malachite, azurite, chrysocolla, neotocite, cuprite, chalcocite, covellite, djurleite, digenite, anilite, and idaite. Cook (1988) described copper-bearing mineral assemblages at the Lakeshore mine in Arizona, as follows: chalcocite zone, brochantite zone, chrysocolla zone, copper wad zone, and goethite zone. These zones do not form distinct horizons but occur throughout the supergene orebodies.

Zoning and Geochemistry

Supergene enrichment processes are important because potential ore metals, such as copper and silver, are concentrated during supergene oxidation, while other metals are removed by the solutions (Hartley and Rice, 2005). At the Lakeshore mine in Arizona, Cook (1988) found chrysocolla mineralization to be in a particular host rock, granodiorite porphyry, with half of the copper in altered phenocrysts of biotite and plagioclase together with clay minerals. There was a complete range of chemical compositions from chrysocolla to copper-bearing clay. Other chrysocolla deposition at Lakeshore was fracture controlled. In the brochantite zone, Cook (1988) found the brochantite to occur principally in fractures and oxidized hypogene quartz veinlets. The goethite zone at Lakeshore occurs as a layer against overlying postmineralization rocks. Although the goethite zone carries 0.3 percent copper, blue and green copper oxide minerals are rare. The cuprite zone occurs as a layer separating brochantite and chrysocolla zones from the chalcocite zone in the deepest part of the supergene orebody.

The geochemistry is reflective of the oxide mineral assemblage; thus, its complexity depends on the complexity of the assemblage mix. For example, in a profile across the goethite, chrysocolla, and chalcocite zones at Lakeshore, Cook (1988) found SiO_2 , Al_2O_3 , MgO , and Na_2O to be depleted in all three zones, but the amount of SiO_2 depletion is a function of the copper content. The silica depletion reflected the amount of chrysocolla present; that is, the greatest depletion was in the goethite zone. The ratio of Fe^{+++} to Fe^{++} increased, but the total iron concentration did not. The CaO was depleted in the goethite zone, but gypsum was present in the chalcocite zone. Copper oxide was enriched across the whole profile. For additional discussion, the reader is referred to Anderson (1982), Alpers and Brimhall (1989), Chávez (2000), and Sillitoe (2005).

Rates

A detailed understanding of the importance of multiple weathering cycles and weathering rates requires careful petrographic study and detailed age information. Although analytical methods such as $^{40}\text{Ar}/^{39}\text{Ar}$ have been applied to the problem, a great deal of care in sampling, sample preparation and analysis, and data interpretation is required to obtain meaningful results. Growth bands indicative of multiple depositional events often can only be documented as a period during which oxidation took place but may not yield rates of weathering (for example, Mote and others, 2001). At El Salvador, Chile, Mote and others (2001) found weathering to have begun within 6 m.y. (about 36–35 Ma) of the end of hypogene mineralization at about 41 Ma. Multiple supergene events were identified (35 Ma, 25 Ma, 14 Ma, 11 Ma), with the youngest episode being the ~11 Ma event. Mote and others (2001) concluded that the supergene events were driven by regional to global climatological events, including global climate change near the Eocene-Oligocene boundary. For additional discussion of rates, the reader is referred to Sillitoe (2005).

Geoenvironmental Features

The following discussions are expanded from Cox and others (1995).

Pre-Mining Baseline Signatures in Soil, Sediment, and Water

Baseline characterization studies of porphyry copper deposits are limited in the literature, particularly in regard to the diversity of climatic settings in which they are found. Therefore, extrapolation of the insights provided by the case studies summarized herein to other deposits should be done with caution. Data from tropical Tanamá, Puerto Rico (Learned and others, 1972), show different metal abundances in the B-horizon of residual soil over magnetite-rich (1.4 wt-percent sulfur) and pyrite-rich (2.5 wt-percent sulfur) copper ore zones. These data illustrate how acid, which is released as pyrite oxidizes, leaches metals from soil. An exception is molybdenum, which is not readily leached in low-pH environments. Chaffee and others (1981) indicated that stream sediment from the area draining the unmined Red Mountain porphyry copper deposit in arid southeastern Arizona contains elevated abundances of molybdenum, lead, and tellurium. The Mineral Butte and Vekol, Arizona, deposits are other arid-climate porphyry-copper deposits for which pre-mining geochemical data are available. Chaffee (1976, table 1, and 1977, table 3) presents ranges, backgrounds, and geochemical anomaly thresholds for manganese, cadmium, copper, cobalt, fluorine, gold, lead, bismuth, mercury, molybdenum, silver, and zinc abundances in bedrock and residual soil samples collected in the vicinity of these deposits. Abundances of most of these elements are elevated in the vicinity of these deposits relative to those characteristic of other geologic environments; these elevated abundances identify appropriate baseline geochemical values that may be useful in setting remediation standards for mines associated with this deposit type in arid climates.

Leybourne and Cameron (2006) documented high-salinity ground waters associated with the unmined Spence deposit in the Atacama Desert, Chile, reaching 10,000 to 55,000 mg/L, with one outlier at 145,000 mg/L. The pH of the ground waters varied widely between 4.7 and 9.2. Dissolved sulfate concentrations in ground waters were between 5,000 and 10,000 mg/L. Plaza-Toledo (2005) found surface waters downstream from unmined porphyry copper deposits in the Cordillera Central of Puerto Rico to reach maximum dissolved concentrations of 110 mg/L sulfate, 0.56 mg/L aluminum, 0.13 mg/L copper, 15.9 mg/L iron, and 0.04 mg/L zinc. The pH was high (7.7–8.6).

Past and Future Mining Methods and Ore Treatment

Most porphyry copper deposits are mined by open-pit methods and, less commonly, by underground methods. For those deposits with associated skarns, underground or open-pit mining methods may be used to mine the skarn deposits. Sulfide ores are generally crushed to a fine-grain size, and ore minerals (chalcopyrite, bornite, molybdenite, and pyrite, if gold-bearing) are separated by conventional flotation methods to form a concentrate. Gold-bearing pyrite concentrates may be roasted onsite to oxidize the pyrite prior to treatment with cyanide to extract the gold. Low-grade and oxide ores may undergo less intensive crushing prior to being treated by solvent-extraction electrowinning (SX–EW) methods with acidic solutions on leach pads to extract copper (Lynch and others, 1994).

Volume of Mine Waste and Tailings

Due to the low percentages of ore minerals typically associated with porphyry copper deposits, the tonnages of flotation tailings and SX–EW leach-pad wastes are essentially identical to the tonnages of the deposits being mined. Volumes of waste rock will depend upon the depth of the deposit, the geometry of the deposit, and the competency of the country rock as it relates

to stripping ratios. On average, about 1.5 tons of waste rocks and overburden must be removed for every ton of ore grade mined (Phillips and Niemuth, 1993).

Mine Waste Characteristics

Mineralogy: In addition to the primary minerals associated with the ores, a number of secondary minerals can form in the mine wastes. These include jarosite, schwertmannite, ferrihydrite, goethite, manganese hydroxides, authigenic clays, covellite, pickeringite, bonattite, chalcantinite, melanterite, and rozenite (Dold and Fontboté, 2001; Hansen and others, 2005).

Acid-Base Accounting

Net neutralization potentials reported in the literature for tailings from several porphyry copper deposits in Chile are net acid, ranging from -101.6 to -18.2 kilograms calcium carbonate per ton ($\text{kg CaCO}_3/\text{t}$) (Dold and Fontboté, 2001). Net neutralizing potentials for hypogene and supergene ores from Morenci, Arizona, range from dominantly net acid to slightly net alkaline (-257.0 to 1.1 $\text{kg CaCO}_3/\text{t}$; Enders and others, 2006).

Metal Mobility Related to Mining in Ground Water and Surface Water

The interaction between a porphyry copper deposit and the environment is illustrated by a stream in the Globe mining district in Arizona that was blocked by mill tailings, causing a lake to form. Water from this lake entered an alluvial aquifer by seepage; the aquifer and a stream to the north were contaminated (Eychaner, 1991; Stollenwerk, 1994; Brown and others, 1998; Lind and others, 1998; Conklin and others, 2001). The most contaminated ground water in the aquifer had a pH of 3.3 and contained about 9,600 mg/L sulfate, 2,800 mg/L iron, 300 mg/L aluminum, and 190 mg/L copper. As the plume traveled north through the aquifer, the concentration of constituents decreased as the plume interacted with carbonate-bearing alluvium and was diluted by uncontaminated water (ground water flowing upward from lower basin fill, water in uncontaminated streams that join the contaminated wash, and surface rainwater).

Enders and others (2006) reported analyses of seeps and springs in the Morenci district of Arizona, which had pH values between 2.6 and 4.6, and specific conductance values between 1,400 and 6,000 $\mu\text{S}/\text{cm}$. Sulfate concentrations were between 550 and 4,300 mg/L, iron between 15 and 420 mg/L, aluminum between 0.48 and 370 mg/L, copper between 0.46 and 960 mg/L, and zinc between 0.8 and 159 mg/L.

Pit Lakes

Pit lakes, particularly in porphyry copper mining districts that are riddled with historical underground mine workings, such as that at Butte, Montana, which has a 140-year mining history, can be problematic. The Berkeley pit lake at Butte contains more than 100 billion liters of pH-2.5 mine water (Gammons and others, 2005; Gammons and Duaiame, 2005). The pit lake has high dissolved solids greater than 7 $\mu\text{S}/\text{cm}$) and elevated concentrations of iron, copper, zinc, and sulfate (Gammons and Duaiame, 2005).

Ecosystem Issues

Acidic mine drainage and associated dissolved metals may pose threats to surrounding aquatic ecosystems depending upon geologic and hydrologic setting and engineering aspects of the mine and the waste piles. Host rocks with higher neutralization potentials, such as carbonate rocks, tend to limit the mobility of metals and related compounds. Hydrologic and climatic settings that are subject to distinct wetting and drying cycles associated with waste piles tend to promote the formation of efflorescent metal sulfate salts that may cause acute toxic effects to aquatic ecosystems if drainage directly enters surface waters. Improperly constructed water-containment structures may allow contaminated mine waters to enter surrounding surface waters.

Acidity will be associated with the pH of the water and the dissolved concentrations of ferrous and ferric iron, aluminum, and manganese; thus, because of incomplete hydrolysis of these elements and incomplete oxidation of dissolved iron and manganese, the pH may provide an inaccurate portrayal of the acidity of mine drainage (and Cravotta, 2005 a, b). Elements or compounds with the greatest likelihood of causing problems for aquatic ecosystems and drinking-water sources for terrestrial organisms include, in alphabetical order, aluminum, arsenic, copper, iron, manganese, sulfate, and zinc.

Human Health Issues

Contaminated ground-water plumes associated with tailings impoundments may pose threats to drinking-water supplies, depending upon hydrologic and geologic setting and engineering aspects of the mine and the waste piles. Host rocks with higher neutralization potentials, such as carbonate rocks, tend to limit the mobility of metals and related compounds. Hydrologic and climatic settings that experience net evaporative loss of water may cause evaporative concentration of solutes that may enter ground water used as drinking-water supplies in the vicinity of mines and waste piles. Improperly constructed water-containment structures may allow contaminated mine waters to enter surrounding ground water. Elements or compounds with the greatest likelihood of causing problems for aquatic ecosystems, and drinking-water sources for terrestrial organisms include, in alphabetical order, aluminum, arsenic, copper, iron, manganese, sulfate, and zinc.

Climate Effects on Geoenvironmental Signatures

The understanding of the effects of various climate regimes on the geoenvironmental signature specific to porphyry copper deposits is limited. In most cases, the intensity of environmental impact associated with sulfide-mineral-bearing mineral deposits is greater in wet climates than in dry climates. Acidity and total metal concentrations in mine drainage in arid environments are several orders of magnitude greater than in more temperate climates because of the concentrating effects of mine-effluent evaporation and the resulting “storage” of metals and acidity in highly soluble metal-sulfate-salt minerals. However, minimal surface-water flow in these areas inhibits generation of significant volumes of highly acidic, metal-enriched drainage. Concentrated release of these stored contaminants to local watersheds may be initiated by precipitation following a dry spell. Present surface and near-surface chemical relations and mineral assemblages may reflect pre-existing (wetter) climate in the Southwestern United States. Dold and Fontboté (2001) investigated the environmental-geochemical characteristics of tailings from Chilean porphyry copper deposits in hyperarid, Mediterranean, and alpine climates. They concluded that climatic variations provide fundamental controls on the mobility of metals. Under precipitation-dominant conditions, divalent metals such as copper, zinc, and manganese are leached in the oxidized zone of mine waste piles, are carried downward, and are sequestered in the piles due to

replacement sulfide minerals in the reduced zone of mine waste piles. Under evaporation-dominant conditions, leached solutes are carried upward due to capillary action and form efflorescent metal-sulfate salts, which are then available for later dissolution during heavy rainstorms or snowmelt events.

Grade and Tonnage Models

Grades and tonnages for porphyry copper deposits are detailed in a complementary open-file report by Singer and others (2008).

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Appendix

It has been recognized for many years that there is variation between porphyry copper deposits regarding the content and concentrations of by-product metals in individual deposits including molybdenum and gold. In the course of constructing grade and tonnage models of porphyry copper deposits by the U.S. Geological Survey, it was found that porphyry copper deposits should be subdivided into three subclasses—copper, copper-molybdenum, and copper-gold—owing to their having statistically different grade and tonnage distributions (see Singer and others, 1986; Singer and Cox, 1986). The first subclass, copper, is referred to as the “general model” because it includes the deposits in the other two subclasses. Cox and Singer (1988) further evaluated the distribution of gold in 55 porphyry copper deposits for which grade and tonnage data were available. Although Cox and Singer (1988) found there to be a continuum in copper, molybdenum, and gold contents between the deposits in their study, they defined three subtypes as follows: copper-gold, copper-gold-molybdenum, and copper-molybdenum. They defined the subtypes on the basis of the gold:molybdenum ratio. Porphyry copper-gold deposits have gold:molybdenum ratios ≥ 30 , copper-gold-molybdenum deposits have gold:molybdenum ratios greater than 3, but less than 30, and copper-molybdenum deposits have gold:molybdenum ratios ≤ 3 .

A number of studies have suggested linkages between the by-product content of porphyry copper deposits and different geological attributes. Kesler (1973) recognized two types of deposits, copper-molybdenum and copper-gold, and concluded that copper-gold deposits are smaller as a group than copper-molybdenum deposits. He also determined that the differences may be explainable by primary compositional differences between mineralizing systems or different chemical responses to distinct geologic features such as the level of granitoid emplacement, permeability of the wall rocks, or wall-rock chemistry. Sillitoe (1979), defining gold-rich as deposits averaging greater than 0.4 gram gold per ton, concluded that high gold contents are not directly related to tectonic setting, host-rock composition, wall-rock composition, deposit age, level of erosion, orebody size, and presence or absence of phyllic alteration. Sillitoe (1979) did note that gold occurs in the potassic alteration zone in deposits with a high magnetite content. He also suggested that copper-poor, gold-rich deposits should exist. Sillitoe (1993) revisited the subject of gold-rich porphyry copper deposits and concluded that gold-rich deposits do not possess any unique features to distinguish them from copper-molybdenum deposits. Sillitoe (1993) found that 80 percent of the 29 deposits he studied have greater magnetite contents than non-gold-rich varieties. He also noted that calc-silicate alteration minerals, including amphibole, pyroxene, and(or) garnet, tend to occur in gold-rich varieties. Both copper-molybdenum and copper-gold deposits occur with calc-alkaline intrusive complexes, but only gold-rich deposits occur in alkalic intrusive complexes. However, those deposits with the most gold occur with calc-alkalic rocks rather than alkalic rock complexes (Sillitoe, 1993).

Table 1. Intrusive rock compositions associated with selected porphyry copper deposits.

Deposit	Intrusive rock types	Reference
Batu Hijau, Indonesia	Hornblende-plagioclase andesite, plagioclase andesite, pyroxene-plagioclase andesite, hornblende diorite, quartz diorite, granodiorite, tonalite, diorite, porphyritic hornblende tonalite, porphyritic dacite, tonalite porphyry	Garwin (2002)
Bingham, Utah, USA	Monzonite, quartz monzonite porphyry, latite porphyry, quartz latite porphyry	Babcock and others (1995)
Cadia, New South Wales, Australia	Monzodiorite, diorite, gabbro, quartz monzonite porphyry, monzonite, latite, trachyte, basalt, basaltic andesite	Holliday and others (2002)
Chuquicamata, Chile	Granodiorite	Ossandón and others (2001)
Copper Mountain, British Columbia, Canada	Diorite, monzonite, syenite, monzodiorite	Stanley and others (1995)
Dexing, Jiangxi, China	Diorite porphyry, granodiorite porphyry, gabbro, diabase	Singer and others (2008)
Dinkidi, Philippines	Diorite, monzonite porphyry, monzonite, quartz monzonite porphyry, syenite	Garrett (1996)
Elatsite, Bulgaria	Diorite porphyry, granodiorite porphyry, monzonite porphyry, quartz diorite porphyry, syenite porphyry, aplite	Singer and others (2008)
El Teniente, Chile	Quartz diorite, tonalite, microdiorite, andesite	Klemm and others (2007) Stern and others (2007)
Escondida, Chile	Quartz monzonite, granodiorite	Padilla-Garza and others (2004)
Grasberg, Indonesia	Andesite, diorite porphyry, monzodiorite porphyry, hornblende monzonite porphyry	MacDonald and Arnold (1994)
Kounrad, Kazakhstan	Diorite porphyry, granodiorite, granodiorite porphyry, quartz diorite porphyry, trondhjemite	Singer and others (2008)
Los Bronces-Rio Blanco, Chile	Quartz monzonite, quartz monzodiorite, granodiorite, hornblende diorite, syenite, aplite, dacite porphyry, quartz latite, tonalite porphyry	Warnaars and others (1985) Deckart and others (2005)
Los Pelambres, Chile	Tonalite, tonalite porphyry	Sillitoe (1973)
Michiquillay, Peru	Quartz monzonite, porphyritic granodiorite	Hollister and Sirvas (1974)
Morenci, Arizona, USA	Diorite porphyry, diorite porphyry, quartz monzonite porphyry, diabase	Moolick and Durek (1966)
Rosia Poieni, Romania	Hornblende andesite, microdiorite	Milu and others (2004)
Sar Cheshmeh, Iran	Granodiorite porphyry, granodiorite, diorite, dacite porphyry, andesite	Waterman and Hamilton (1975) Hezarkhani (2006)

Table 2. Major and minor axes of extent of altered rock, orebodies, and sulfide-bearing rock in kilometers together with the total area of altered rock, ore, and sulfide-bearing area, in square kilometers for selected deposits, worldwide. The shortest dimension (minor axis) is measured by the distance between parallel lines that just touch the object. The major axis is measured to the ends of the object on a line perpendicular to the short axis. Areas are calculated assuming an elliptical object in plan view. The deposits are organized alphabetically by country. Dashed lines indicate no data available. (After Singer and others, 2008).

Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Agua Rica, Argentina	4.4	3.8	13.1	13.1	--	--	--	--	--
Bajo de la Alumbrera, Argentina	2.5	1.75	3.4	1.1	0.6	0.52	--	--	--
Los Azules, Argentina	5	8	31.4	3.3	0.9	2.3	--	--	--
San Jorge, Armenia	--	--	--	0.8	0.4	0.25	--	--	--
Agarak, Armenia	--	--	--	0.7	0.25	0.14	--	--	--
Cadia Hill/Ridgeway, Australia	7	3	16.5	1.0	0.35	0.27	3.5	1.6	4.4
Coalstoun, Australia	1.8	0.8	1.1	0.45	0.2	0.07	1.2	0.7	0.66
Coppin Gap, Australia	--	--	--	1	0.5	0.39	--	--	--
Endeavour, Australia	6.5	3.7	19	1.7	0.96	1.3	--	--	--
Mount Cannindah, Australia	--	--	--	1.7	0.96	1.3	--	--	--
Peak Hill, Australia	3.3	0.6	1.6	--	--	--	--	--	--
Yeoval, Azerbaijan	--	--	--	0.5	0.3	0.12	--	--	--
Yeppon, Azerbaijan	--	--	--	0.9	0.35	0.25	--	--	--
Chapada, Brazil	--	--	--	2.7	0.7	1.5	--	--	--
Assarel, Bulgaria	2.8	1	2.2	--	--	--	--	--	--
Elatsite, Bulgaria	--	--	--	1	0.6	0.47	--	--	--
Medet, Bulgaria	--	--	--	1	0.4	0.31	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Ajax, Canada	--	--	--	0.42	0.15	0.05	--	--	--
Bell Copper, Canada	3.2	2.6	6.5	0.63	0.43	0.21	2.1	2	3.3
Berg, Canada	--	--	--	1.8	1.7	2.4	2.4	1.9	3.5
Bethlehem, Canada	--	--	--	--	--	--	1.8	1.4	1.9
Big Onion, Canada	--	--	--	1.9	0.35	0.65	--	--	--
Brenda, Canada	--	--	--	0.92	0.48	0.35	3.2	2.4	6
Casino, Canada	--	--	--	1.8	0.76	1.1	2.4	1.4	2.7
Copper Mountain, Canada	--	--	--	4.3	1.1	3.7	6	3	14
Don Rouyn, Canada	--	--	--	0.5	0.2	0.08	--	--	--
Dorothy, Canada	2.3	1.8	3.3	0.5	0.3	0.1	1.6	1.2	1.5
Fish Lake, Canada	--	--	--	1.6	0.8	0.97	3.6	2.5	7.1
Galore Creek, Canada	--	--	--	1.9	0.5	0.75	5	4	16
Gambier Island, Canada	--	--	--	1.1	0.6	0.52	--	--	--
Gaspé, Canada	9	5	7.5	0.6	0.3	0.14	2.5	1.1	2.2
Granisle, Canada	3.2	2.2	5.6	0.7	0.3	0.16	--	--	--
Highmont, Canada	--	--	--	1.1	0.73	0.64	--	--	--
Huckleberry, Canada	5	2.5	9.8	1	0.3	0.24	3.9	1.5	1.4
Hushamu, Canada	6	3	5	1.8	0.9	1.3	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Island Copper, Canada	5.2	1.1	4.9	2	0.52	0.31	2.2	1	1.7
Kemess North, Canada	6.7	2.6	14	1.2	0.6	0.57	--	--	--
Kemess South, Canada	--	--	--	1.7	0.75	1	--	--	--
Lornex, Canada	--	--	--	1.9	1.1	1.3	--	--	--
Louise Lake, Canada	4	1	3.1	0.85	0.35	0.23	--	--	--
Maggie, Canada	--	--	--	1.1	0.33	0.26	3.7	1.5	4.4
McIntyre, Canada	2.5	0.5	0.98	0.5	0.13	0.05	--	--	--
Morrison, Canada	--	--	--	1.5	0.7	0.82	2.8	2.3	5.1
Ox Lake, Canada	1.3	1.3	1.3	0.7	0.5	0.27	1	0.8	0.63
Poison Mountain, Canada	--	--	--	1.5	0.8	0.94	1.7	1	1.3
Red Chris, Canada	5	1.5	5.9	1.4	0.5	0.55	--	--	--
Schaft Creek, Canada	3.1	1.3	3.1	2.5	1.1	2.2	--	--	--
Sulphurets, Canada	6	3	14	1.8	0.9	1.3	4	1.8	5.6
Valley, Canada	--	--	--	1.8	1.2	1.7	6.5	4	20.4
Willa, Canada	--	--	--	0.5	0.15	0.06	--	--	--
Andacollo, Chile	5	3.8	15	1.5	1.3	1.5	5	2.5	9.8
Chuquicamata, Chile	--	--	--	12	3	28	--	--	--
Collahuasi, Chile	7	4.7	26	1.9	1.2	1.8	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
El Abra, Chile	--	--	--	1.7	1.2	1.6	4.4	2.3	8
El Salvador, Chile	8	5	31	1.8	1	1.4	3	3	7.1
El Teniente, Chile	--	--	--	2.7	2	4.2	4	3.5	11
Escondida, Chile	11	9	78	4.5	1.6	5.7	--	--	--
Lomas Bayas, Chile	--	--	--	1.6	1.3	1.6	--	--	--
Los Bronces/Rio Blanco, Chile	--	--	--	5	2.8	11	11	5	43
Los Pelambres, Chile	6.5	4.6	24	1.6	0.5	0.6	6.5	2.5	13
Potrerrillos, Chile	--	--	--	1.5	0.76	0.91	6	4	19
Chengmenshan, China	--	--	--	0.8	0.6	0.4	--	--	--
Dexing, China	7.5	2	12	1.1	0.9	0.74	--	--	--
Kalatage, China	--	--	--	0.6	0.04	0.02	--	--	--
Malasongduo, China	--	--	--	0.95	0.8	0.6	--	--	--
Mangzhong, China	--	--	--	1.1	0.25	0.22	--	--	--
Nanmu, China	2.8	0.9	2	1.2	0.3	0.3	--	--	--
Tinggong, China	--	--	--	1.4	0.6	0.7	--	--	--
Tuwu, China	3	0.4	0.9	2.3	0.4	0.4	--	--	--
Xietongmen, China	--	--	--	0.9	0.3	0.2	--	--	--
Xifanping, China	1.5	0.8	0.9	0.7	0.4	0.2	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Yandong, China	--	--	--	0.9	0.2	0.14	--	--	--
Yulong, China	2	1.5	2.4	1	0.8	0.6	--	--	--
Zijinshan, China	--	--	--	0.83	0.5	0.32	--	--	--
El Cobre, Cuba	--	--	--	1.8	0.7	0.99	--	--	--
Chaucha, Ecuador	--	--	--	3.3	0.95	2.5	6.8	2.9	16
Cumay, Ecuador	--	--	--	1.5	0.75	0.88	3.5	1.5	4.1
Gaby-Papa Grande, Ecuador	--	--	--	0.9	0.1	0.07	4.4	1.3	4.5
Junin, Ecuador	2.5	2	3.9	0.4	0.4	0.13	--	--	--
Panantza, Ecuador	--	--	--	0.8	0.2	0.04	--	--	--
San Carlos, Ecuador	--	--	--	0.6	0.6	0.28	1.9	1.2	1.8
Namosi, Fiji	6	4.8	23	--	--	--	--	--	--
Recsk, Hungary	--	--	--	2	1	1.6	3	1.4	3.4
Malanjkhanda, India	--	--	--	2.6	0.7	0.72	--	--	--
Batu Hijau, Indonesia	6.5	3.5	18	1.2	0.6	0.57	3.5	2.5	6.9
Grasberg, Indonesia	6	3	14	1	0.8	0.6	2	1.7	2.7
Kaputusan, Indonesia	1.8	0.9	1.3	1	0.3	0.2	--	--	--
Kale Kafi, Iran	--	--	--	1.4	0.7	0.77	--	--	--
Meiduk, Iran	--	--	--	0.54	0.37	0.2	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Sar Cheshmeh, Iran	--	--	--	2.5	1	2	--	--	--
Sungun, Iran	--	--	--	2.3	1.5	2.7	--	--	--
Aktogai, Kazakhstan	9.2	1	7.2	--	--	--	--	--	--
Boshchekul, Kazakhstan	6.3	1.5	7.4	3	0.5	1.2	--	--	--
Kounrad, Kazakhstan	--	--	--	1.2	0.8	0.75	--	--	--
Taldy-Bulak, Kirgystan	--	--	--	1.2	0.8	0.75	--	--	--
Kadiica, Macedonia	--	--	--	1	0.6	0.5	--	--	--
Mamut, Malaysia	--	--	--	0.7	0.5	0.3	2.2	1.3	2.3
Nungkok, Malaysia	--	--	--	0.4	0.15	0.05	--	--	--
Cananea, Mexico	--	--	--	3	1.5	3.5	7	4	22
El Arco, Mexico	3.5	3.5	9.6	1.5	1.4	1.7	--	--	--
La Caridad, Mexico	4.5	4	14	2.4	1.7	3.2	--	--	--
Milpillas, Mexico	3.5	3.2	8.8	1.5	1.2	1.4	--	--	--
Piedras Verdes, Mexico	8	3	19	3	0.45	1.1	--	--	--
Avdartolgoi, Mongolia	--	--	--	0.75	0.45	0.27	--	--	--
Bayan Uul, Mongolia	5	3	11.8	--	--	--	--	--	--
Erdenet, Mongolia	--	--	--	2.8	0.8	1.8	4.5	1.5	5.3
Khongor, Mongolia	--	--	--	1.7	0.33	0.44	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Oyu Tolgoi, Mongolia	--	--	--	6.5	0.6	3.1	--	--	--
Oyuut Ulaar Ovoo, Mongolia	--	--	--	0.7	0.25	0.14	--	--	--
Tsagaan-Suvarga, Mongolia	--	--	--	0.9	0.2	0.14	10	2	16
Haib, Namibia	13	3	31	3.1	1.3	3.2	13	3	31
Saindak, Pakistan	6.8	2	11	1.9	0.4	0.6	3.5	1.7	4.7
Cerro Colorado, Panama	2.5	1.9	3.7	1.8	1.1	1.6	--	--	--
Frieda River, Papua-New Guinea	10	4	47	--	--	--	--	--	--
Ok Tedi, Papua-New Guinea	3.6	1.8	5.1	1.8	0.77	1.1	--	--	--
Panguna, Papua-New Guinea	6	5.3	25	3	2	4.7	--	--	--
Yandera, Papua-New Guinea	8	3	19	1.6	0.25	0.3	--	--	--
Cuajone, Peru	--	--	--	2	1	1.6	5	4	16
Magistral, Peru	--	--	--	1.2	0.13	0.12	--	--	--
Michiquillay, Peru	2.3	2	3.6	0.81	0.75	0.48	2	2	3.1
Minas Conga, Peru	--	--	--	1.3	0.6	0.59	--	--	--
Pashpap, Peru	1.5	1.5	1.8	1.2	0.3	0.3	--	--	--
Quellaveco, Peru	--	--	--	2	1.1	1.7	3	1.5	3.5
Toquepala, Peru	--	--	--	2.1	1.8	3	--	--	--
Toromocho, Peru	--	--	--	1.4	1	1.1	--	--	--

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Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Atlas, Philippines	--	--	--	3.1	0.75	1.8	--	--	60
Basay, Philippines	--	--	--	1.8	0.25	0.34	--	--	4.5
Boneng Lobo, Philippines	--	--	--	0.4	0.3	0.09	--	--	1.7
Boyongan, Philippines	--	--	--	1.2	0.5	0.47	--	--	--
Dinkidi, Philippines	--	--	--	0.45	0.15	0.05	--	--	--
Dizon, Philippines	--	--	--	0.55	0.4	0.17	--	--	8
Far Southeast, Philippines	4.5	2	7.1	--	--	--	1	0.7	0.5
Guinaoang-Tirad, Philippines	2.3	2.1	3.8	1.1	0.5	0.43	--	--	--
Hinobaan, Philippines	--	--	--	1.5	0.45	0.53	--	--	--
Marcopper, Philippines	3.4	2.5	6.7	1.7	0.64	0.85	2.7	2	4.2
Matanlang, Philippines	--	--	--	0.8	0.6	0.5	--	--	--
San Antonio-Philex, Philippines	--	--	--	0.26	0.2	0.04	--	--	0.5
San Fabian, Philippines	--	--	--	1.3	0.3	0.3	--	--	1
Santo Tomas II, Philippines	--	--	--	0.57	0.46	0.21	--	--	1.3
Sipalay, Philippines	--	--	--	1.5	0.6	0.7	--	--	--
Taysan, Philippines	2.5	1.513	--	0.5	0.4	--	--	--	--
Bolcana, Romania	3	1.3	3.1	0.54	0.4	0.17	--	--	--

Table 2. Major and minor axes of extent of altered rock, orebodies, and sulfide-bearing rock in kilometers together with the total area of altered rock, ore, and sulfide-bearing area, in square kilometers for selected deposits, worldwide. The shortest dimension (minor axis) is measured by the distance between parallel lines that just touch the object. The major axis is measured to the ends of the object on a line perpendicular to the short axis. Areas are calculated assuming an elliptical object in plan view. The deposits are organized alphabetically by country. Dashed lines indicate no data available. (After Singer and others, 2008).—Continued

Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Talagiu, Romania	5.5	2.8	12	2.1	0.8	1.3	--	--	--
Birgildinskoe, Russia	--	--	--	0.6	0.1	0.05	1.2	0.3	0.28
Nakhodka, Russia	--	--	--	2.8	1.1	2.4	--	--	--
Peschanka, Russia	--	--	--	4.4	0.7	2.4	--	--	--
Salavat, Russia	--	--	--	1.3	0.4	0.41	--	--	--
Bor, Serbia	--	--	--	2	2.4	1.5	--	--	--
Majdanpek, Serbia	--	--	--	4	0.3	0.94	--	--	--
Rudnitsa, Serbia	1.2	0.8	0.75	0.4	0.3	0.09	--	--	--
Veliki Krivelj, Serbia	--	--	--	1.5	0.7	0.8	--	--	--
Chimei, Taiwan	--	--	--	0.6	0.1	0.05	2.5	2	3.9
Derekoy, Turkey	5	0.75	2.9	2.9	0.7	1.6	--	--	--
Bagdad, USA	--	--	--	1.5	0.66	0.79	--	--	--
Bingham, USA	--	--	--	2.4	1.3	2.4	6.5	4.8	25
Bisbee, USA	2.7	2.7	5.9	0.61	0.61	0.29	--	--	--
Castle Dome, USA	2.7	1.5	3.3	0.91	0.46	0.33	--	--	2.6
Christmas, USA	2	1.7	2.7	1.5	0.9	1.1	--	--	--
Copper Basin, USA	--	--	--	1.4	1.4	1.5	--	--	--
Copper Creek, USA	--	--	--	1.7	0.58	0.77	--	--	7.5

Table 2. Major and minor axes of extent of altered rock, orebodies, and sulfide-bearing rock in kilometers together with the total area of altered rock, ore, and sulfide-bearing area, in square kilometers for selected deposits, worldwide. The shortest dimension (minor axis) is measured by the distance between parallel lines that just touch the object. The major axis is measured to the ends of the object on a line perpendicular to the short axis. Areas are calculated assuming an elliptical object in plan view. The deposits are organized alphabetically by country. Dashed lines indicate no data available. (After Singer and others, 2008).—Continued

Deposit	Alteration (major axis)	Alteration (minor axis)	Altered area	Ore (major axis)	Ore (minor axis)	Ore area	Sulfides (major axis)	Sulfides (minor axis)	Sulfide area
Copper Flat, USA	--	--	--	0.8	0.5	0.3	--	--	--
Ely, USA	12	1.2	11	7.3	0.91	5.2	--	--	--
Ithaca Peak, USA	4.1	3.7	12	1	0.67	0.55	--	--	--
Mission-Pima, USA	5.5	3.2	14	2.4	1.6	3	5.5	3.2	14
Morenci, USA	--	--	--	4	1.8	5.7	--	--	49
Pebble Copper, USA	--	--	--	1.6	0.8	1	--	--	89
Pyramid, USA	--	--	--	1	0.35	0.28	3.7	3	8.7
Ray, USA	12	8.5	82	3	0.9	2.1	--	--	9.2
Red Mountain, USA	3.5	3	8.2	1.7	0.3	0.4	--	--	--
San Manuel-Kalamazoo, USA	--	--	--	4.6	0.95	3.4	--	--	6.6
Santa Rita, USA	--	--	--	2	1.3	2	--	--	--
Sierrita-Esperanza, USA	--	--	--	2.2	0.89	1.6	--	--	--
Silver Bell, USA	10	3.5	28	2.5	1.3	2.6	4.8	2.4	9.1
Tyrone, USA	--	--	--	2.8	2	4.4	--	--	--

Table 3. Sulfide, supergene, and alteration mineralogy in a representative selection of porphyry-style copper deposits, worldwide. Minerals shown are as abstracted from the references cited and may not constitute a complete listing of minerals that occur in the deposit. Deposits are in alphabetical order by deposit name.

Deposit	Primary sulfide mineralogy	Supergene mineralogy	Primary alteration mineralogy	References
Ajo, Arizona, USA	Chalcopyrite, bornite, pyrite, molybdenite, sphalerite, tetrahedrite, tennantite	Chalcocite, cuprite, chrysocolla, shattuckite, malachite, azurite, native copper, opal, stibiconite, melaconite, copper pitch, brochantite, gypsum, hematite, goethite, psilomelane, beidellite, nontronite, alunite, jarosite	Quartz, orthoclase, sericite, magnetite, ilmenite, anhydrite, rutile, apatite, albite, chlorite, epidote, clay, specularite, dolomite, ankerite	Gilluly (1946) Dixon (1966)
Almalyk (Kal'makyr), Uzbekistan	Chalcopyrite, pyrite, molybdenite, bornite, tetrahedrite, galena, sphalerite	Chalcocite, cuprite, covellite, malachite, azurite, chrysocolla, ehrlite, libethenite	Quartz, K-feldspar, biotite, sericite, magnetite, anhydrite, calcite	Golovanov and others (2005)
Andacollo, Chile	Chalcopyrite, pyrite, bornite, hematite, molybdenite, pyrrhotite	Chalcocite, azurite, malachite, chrysocolla, cubanite	Quartz, K-feldspar, biotite, anhydrite, magnetite, sericite, rutile, chlorite, epidote, calcite	Reyes (1991) Singer and others (2008)
Bajo de la Alumbrera, Argentina	Chalcopyrite, pyrite	Chalcocite, digenite, covellite with goethite, jarosite	Quartz, orthoclase, biotite, magnetite, anhydrite, hematite, sericite, calcite	Guilbert (1995) Singer and others (2008)
Batu Hijau, Sumbawa, Indonesia	Chalcopyrite, bornite, pyrite, chalcocite, digenite, native gold, sphalerite, galena, tennantite	Chalcocite, covellite, cuprite	Quartz, biotite, albite-oligoclase, magnetite, sericite, actinolite, epidote, chlorite, calcite, specularite, kaolinite, alunite, pyrophyllite, tourmaline	Meldrum and others (1994) Garwin (2002)
Bingham, Utah, USA	Chalcopyrite, bornite, pyrite, molybdenite, chalcocite, digenite, covellite, enargite, famatinite, galena, sphalerite, tetrahedrite	Chalcocite, covellite, anglesite, azurite, cerussite, chalcantite, cuprite, goslarite, malachite, melanterite, native copper, native silver, native gold, tenorite	Quartz, K-feldspar, biotite, apatite, actinolite, sericite, kaolinite, montmorillonite, chlorite, epidote, calcite	Peters and others (1966) Babcock and others (1995)
Cadia, New South Wales, Australia	Chalcopyrite, bornite, molybdenite, pyrite, native gold, covellite, sphalerite, galena, tetrahedrite	Chalcocite, digenite, covellite	Quartz, orthoclase, biotite, magnetite, sericite, chlorite, albite, actinolite, epidote, calcite, tourmaline, hematite, fluorite	Holliday and others (2002) Wilson and others (2007) Singer and others (2008)
Campana Mahuida, Argentina	Chalcopyrite, bornite, molybdenite, pyrite, pyrrhotite, native gold	Chalcocite, covellite, chrysocolla, tenorite	Biotite, quartz, orthoclase, magnetite, sericite, albite, tourmaline, chlorite, epidote, calcite, illite, smectite	Franchini, and others (2007) Singer and others (2008)
Cananea, Sonora, Mexico	Chalcopyrite, pyrite, bornite, molybdenite, sphalerite, galena, enargite, luzonite	Chalcocite, digenite, covellite	Quartz, K-feldspar, biotite, sericite, tourmaline, magnetite, actinolite, chlorite, carbonate, epidote, alunite	Velasco (1966) Singer and others (2008)

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Deposit	Primary sulfide mineralogy	Supergene mineralogy	Primary alteration mineralogy	References
Castle Dome, Arizona, USA	Chalcopyrite, pyrite, molybdenite, sphalerite, galena	Chalcocite, covellite, azurite, malachite, native copper, cuprite, chalcantite, native silver, turquoise	Quartz, K-feldspar, biotite, sericite, anhydrite, chlorite, epidote, clinozoisite, calcite, tourmaline, fluorite, barite	Peterson and others (1946) Singer and others (2008)
Chuquicamata, Chile	Chalcopyrite, bornite, pyrite, molybdenite, digenite, enargite, covellite, tennantite, sphalerite	Chalcocite, covellite, djurleite, digenite, antlerite, brochantite, atacamite, chrysocolla, copper pitch, chenevixite, chalcantite, krohnkite, natrochalcite	Quartz, K-feldspar, biotite, anhydrite, sericite, magnetite, albite, chlorite, epidote, specularite	Ossandón C. and others (2001) Sillitoe (2005)
Copper Cities, Arizona, USA	Chalcopyrite, pyrite, molybdenite, sphalerite, galena	Chalcocite, covellite, malachite, azurite, turquoise, metatorbernite with montmorillonite, sericite	Quartz, K-feldspar, biotite, epidote, calcite, clinozoisite	Simmons and Fowells (1966)
Dexing, China	Chalcopyrite, bornite, pyrite, molybdenite, sphalerite, tennantite	Chalcocite, cubanite, cuprite, tenorite, djurleite, malachite	Quartz, K-feldspar, biotite, sericite, anhydrite, magnetite, apatite, chlorite, albite, calcite, hematite, barite, tourmaline	Li and Sasaki (2007) Singer and others (2008)
El Abra, Chile	Chalcopyrite, bornite, pyrite, enargite, chalcocite	Chrysocolla, atacamite, malachite, antlerite, brochantite, copper pitch, cuprite, tenorite, native copper, chalcocite with kaolinite, sericite, montmorillonite, nontronite	Quartz, K-feldspar, biotite, magnetite, ilmenite, perthite, rutile, apatite, anhydrite, sericite, pyrophyllite, kaolinite, chlorite, epidote, albite, specularite, zeolites, carbonates, tourmaline	Ambrus (1977)
El Salvador, Chile	Chalcopyrite, bornite, pyrite, molybdenite, enargite, tennantite, galena, sphalerite	Chalcocite, digenite, covellite, brochantite, chrysocolla	Quartz, K-feldspar, biotite, sericite, Na-plagioclase, anhydrite, magnetite, actinolite, chlorite, rutile, tourmaline, kaolinite, andalusite, corundum, diaspore, pyrophyllite, svenbergite, woodhouseite, zunyite, calcite, epidote, hematite	Gustafson and Hunt (1975) Sillitoe (2005) Singer and others (2008)
El Teniente, Chile	Chalcopyrite, bornite, pyrite, molybdenite, tennantite-tetrahedrite, enargite, galena, sphalerite	Chrysocolla, malachite, brochantite, antlerite, cuprite, native copper, azurite, tenorite, olivenite, chalcophyllite	Quartz, K-feldspar, biotite, anhydrite, magnetite, albite, tourmaline, chlorite, carbonate, epidote, specularite, rutile, apatite	Camus (1975) Sillitoe (2005) Klemm and others (2007)
Ely, Nevada, USA	Chalcopyrite, bornite, pyrite, molybdenite	Chalcocite, covellite	Quartz, K-feldspar, biotite, sericite, magnetite, specularite	Bauer and others (1966)

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Deposit	Primary sulfide mineralogy	Supergene mineralogy	Primary alteration mineralogy	References
Escondida, Chile	Chalcopyrite, pyrite, molybdenite, enargite, covellite, sphalerite, galena, tennantite	Brochantite, antlerite, chrysocolla, neotocite, tenorite, chalcantite, cuprite, native copper, turquoise, libethenite, malachite	Quartz, K-feldspar, biotite, anhydrite, sericite, magnetite, chlorite, epidote, calcite, alunite, pyrophyllite, diaspore, svenbergite	Alpers and Brimhall (1989) Padilla-Garza and others (2004) Sillitoe (2005)
Esperanza, Chile	Chalcopyrite, bornite, pyrite, molybdenite, chalcocite, native gold	Chalcocite, atacamite, brochantite, chrysocolla	Quartz, K-feldspar, biotite, magnetite, anhydrite, chlorite, apatite, rutile, tourmaline, sericite, illite, epidote, calcite	Perrelló and others (2004) Singer and others (2008)
Far South East, Philippines	Chalcopyrite, pyrite, bornite, molybdenite, sphalerite, tetrahedrite, tennantite, enargite, luzonite, marcasite		Quartz, biotite, magnetite, anhydrite, chlorite, tourmaline, alunite, pyrophyllite, diaspore, barite	Singer and others (2008)
Goonumbla, New South Wales, Australia	Chalcopyrite, bornite, pyrite, native gold, chalcocite, tetrahedrite	Malachite, azurite, chrysocolla, native copper, chalcocite	Quartz, K-feldspar, biotite, sericite, magnetite, anhydrite, carbonates, amphibole, chlorite, epidote	Jones (1985)
Grasberg, Indonesia	Chalcopyrite, bornite, pyrite, digenite, covellite	Chalcopyrite with sericite, kaolinite, montmorillonite	Quartz, biotite, K-feldspar, magnetite, anhydrite, sericite, albite, actinolite, kaolinite, chlorite, epidote, hematite	MacDonald and Arnold (1994)
Kingking, Philippines	Chalcopyrite, bornite, native gold, calaverite	Chalcocite	Quartz, K-feldspar, biotite, sericite, anhydrite, magnetite, chlorite, epidote, kaolinite	Suerte and others (2007) Singer and others (2008)
Los Pelambres, Chile	Chalcopyrite, bornite, pyrite, molybdenite	Malachite, azurite, chrysocolla, chalcantite, brochantite, gypsum, chalcocite, djurleite	Quartz, K-feldspar, biotite, anhydrite, sericite, calcite, apatite, chlorite, albite-oligoclase, tourmaline, rutile, magnetite, tourmaline, kaolinite, chlorite, epidote, specularite	Sillitoe (1973)
Michiquillay, Peru	Chalcopyrite, pyrite, molybdenite, sphalerite, galena, enargite, marcasite, arsenopyrite	Chalcocite, covellite	Quartz, K-feldspar, biotite, sericite, magnetite, anhydrite, chlorite, tourmaline, carbonate, alunite, kaolinite, barite, pyrophyllite	Hollister and Sirvas (1974) Singer and others (2008)
Miduk, Iran	Chalcopyrite, pyrite, bornite, molybdenite, sphalerite, galena, tetrahedrite	Chalcocite, covellite, azurite, malachite, chalcantite, chrysocolla, turquoise	Quartz, K-feldspar, biotite, sericite, magnetite, anhydrite, clay, rutile, magnetite, hematite, chlorite, epidote, calcite	Taghipour and others (2008) Singer and others (2008)

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Deposit	Primary sulfide mineralogy	Supergene mineralogy	Primary alteration mineralogy	References
Morenci, Arizona, USA	Chalcopyrite, pyrite, molybdenite, sphalerite, galena, native gold	Chalcocite, covellite, native copper, cuprite, turquoise, chrysocolla, malachite, tenorite, brochantite, azurite chalcantite, copiopite	Quartz, K-feldspar, biotite, sericite, chlorite	Moolick and Durek (1966)
Ok Tedi, Papua-New Guinea	Chalcopyrite, bornite, pyrite, marcasite, molybdenite	Chalcocite, digenite, djurleite, covellite, copper pitch, turquoise, chrysocolla	Quartz, biotite, K-feldspar, rutile, magnetite, hematite	Bamford (1972) Howell and others (1978)
Oyu Tolgoi, Mongolia	Chalcopyrite, bornite, pyrite, chalcocite, enargite, tennantite	Chalcocite, covellite, malachite, chrysocolla, arsenosulvanite, native copper, tenorite	Quartz, biotite, albite, magnetite, muscovite, K-feldspar, actinolite, illite, chlorite, siderite, fluorite, calcite, dolomite, alunite, pyrophyllite, diaspore, zunyite, topaz, corundum, andalusite, kaolinite, dickite	Khashgerel and others (2006) Singer and others (2008)
Panguna, Bougainville, Papua New Guinea	Chalcopyrite, pyrite, bornite, molybdenite, digenite, galena, sphalerite, bravoite, tennantite, stannite, pyrrhotite	Chalcocite, covellite	Quartz, biotite, K-feldspar, magnetite, anhydrite, hematite, rutile, albite, sericite, chlorite, epidote, calcite, actinolite, kaolinite	Eastoe (1978) Ford (1978) Singer and others (2008)
Pebble, Alaska, USA	Chalcopyrite, pyrite, bornite, molybdenite, native gold, galena, sphalerite	Covellite, digenite, chalcocite	Quartz, orthoclase, biotite, sericite, anhydrite, apatite, rutile, albite, carbonate	Bouley and others (1995)
Portrerillos, Chile	Chalcopyrite, pyrite, bornite, arsenopyrite, sphalerite, stibnite, galena, tetrahedrite	Chalcocite, covellite, brochantite, copper pitch, chrysocolla, azurite, cuprite, native copper, malachite, libethenite, native silver	Quartz, K-feldspar, biotite, magnetite, anhydrite, carbonate, alunite, pyrophyllite, diaspore, kaolinite, woodhouseite, zunyite	Sillitoe (2005) Singer and others (2008)
Quebrada Blanca, Chile	Chalcopyrite, bornite, pyrite, molybdenite	Brochantite, chrysocolla, cuprite, native copper, chalcocite	Quartz, K-feldspar, biotite, chlorite, epidote, tourmaline	Sillitoe (2005) Singer and others (2008)
Quellaveco, Peru	Chalcopyrite, pyrite, molybdenite, enargite, sphalerite, galena	Chalcocite, covellite, brochantite, chrysocolla, malachite, copper pitch, cubanite, cuprite, native copper, turquoise	Quartz, sericite, epidote	Sillitoe (2005) Singer and others (2008)
Ray, Arizona, USA	Chalcopyrite, pyrite, bornite, molybdenite, galena, sphalerite	Chalcocite, covellite, native copper, cuprite, chalcotrichite, native silver, chrysocolla	Quartz, biotite, orthoclase, sericite, kaolinite, montmorillonite, chlorite, epidote, albite, carbonate	Metz and Rose (1966)

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Deposit	Primary sulfide mineralogy	Supergene mineralogy	Primary alteration mineralogy	References
Río Blanco-Los Bronces, Chile	Chalcopyrite, bornite, molybdenite, pyrite, luzonite, enargite, tennantite, valleriite, sphalerite, galena, arsenopyrite	Antlerite, brochantite, chalcantinite, copper pitch, neotocite, chrysocolla, chalcocite, covellite, native copper, cuprite	Quartz, biotite, K-feldspar, anhydrite, actinolite, magnetite, sericite, plagioclase, tourmaline, chlorite, hematite, carbonates, epidote, alunite, pyrophyllite, barite	Warnaars and others (1985) Vargas R. and others (1999) Deckart and others (2005) Sillitoe (2005)
Rosia Poieni, Romania	Chalcopyrite, pyrite, enargite, luzonite	Scorodite, goethite, alunite	Quartz, biotite, K-feldspar, perthite, rutile, anhydrite, magnetite, albite, illite, smectite, kaolinite, chlorite, epidote, carbonate, alunite, dickite, pyrophyllite, diaspore, zunyite, minamiite	Milu and others (2004)
San Manuel-Kalamazoo, Arizona, USA	Chalcopyrite, pyrite, bornite, molybdenite	Chalcocite, covellite, bornite, chrysocolla, cuprite, malachite, native copper, atacamite, plancheite	Quartz, K-feldspar, biotite, magnetite, sericite, anhydrite, chlorite, epidote, calcite	Thomas (1966) Sandbak and Alexander (1995)
Santa Rita, New Mexico, USA	Chalcopyrite, bornite, pyrite, molybdenite, marcasite, pyrrhotite, sphalerite, galena	Chalcocite, covellite, native copper, chrysocolla, cuprite, malachite, azurite, turquoise, melaconite, libethenite	Quartz, orthoclase, biotite, epidote, sericite, chlorite	Rose and Baltosser (1966)
Sar Cheshmeh, Iran	Chalcopyrite, bornite, pyrite, molybdenite	Chalcocite, digenite, covellite	Quartz, K-feldspar, biotite, anhydrite, albite, sericite, kaolinite, illite, chlorite, epidote	Waterman and Hamilton (1975) Hezarkhani (2006)
Silver Bell, Arizona, USA	Chalcopyrite, pyrite, molybdenite	Chalcocite, azurite, malachite, brochantite, chalcantinite, cuprite, chrysocolla, native copper, tenorite	Quartz, biotite, sericite, actinolite, chlorite	Richard and Courtright (1966) Singer and others (2008)
Tongshankou, China	Chalcopyrite, pyrite, bornite, sphalerite, tetrahedrite, molybdenite		Quartz, orthoclase, biotite, quartz, magnetite, hematite, sericite, calcite	Li and others (2008)
Toquepala, Peru	Chalcopyrite, pyrite, molybdenite, bornite, tennantite, cobaltite, mackinawite, enargite, sphalerite, galena	Chalcocite, digenite, native copper	Tourmaline, quartz, K-feldspar, albite, biotite, sericite, anhydrite, magnetite, apatite, rutile, chlorite, actinolite, andalusite, siderite, pyrophyllite	Zweng and Clark (1995) Singer and others (2008)
Valley Copper, British Columbia, Canada	Chalcopyrite, bornite, molybdenite, pyrite, enargite, tetrahedrite, sphalerite, galena, scheelite	Chalcocite, digenite, covellite, azurite, malachite, native copper, cuprite, powellite, tenorite	Quartz, K-feldspar, sericite, biotite, anhydrite, magnetite, calcite, hematite, kaolinite, chlorite, epidote, fluorite	Osatenko and Jones (1976) Singer and others (2008)

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Deposit	Primary sulfide mineralogy	Supergene mineralogy	Primary alteration mineralogy	References
Wasoi, Namosi District, Fiji	Chalcopyrite, bornite, pyrite, molybdenite, native gold, galena, sphalerite	Chalcocite, covellite, chrysocolla, native copper, cuprite, malachite, neotocite	Quartz, biotite, magnetite, sericite, chlorite, epidote, carbonate	Imai and others (2007) Singer and others (2008)

Table 4. Alteration-mineral assemblages in the Copper Mountain (British Columbia, Canada) and El Teniente (Chile) porphyry copper deposits.

Pervasive style	Minerals
Copper Mountain, British Columbia, Canada	
Sodic variety	Oligoclase/albite after plagioclase; minor amounts of epidote, diopside, and calcite after ferromagnesian silicates; magnetite destroyed; pyrite only sulfide, generally rare; fine grained where alteration very intense; may be overprinted by potassic variety alteration assemblage <i>Summary:</i> Albite + epidote + (chlorite) +(calcite) ± diopside
Potassic variety	K-feldspar after plagioclase; biotite, epidote, and calcite after ferromagnesian silicates; magnetite present; minor sulfides, predominantly pyrite <i>Summary:</i> K-feldspar + biotite + magnetite + epidote ± chlorite
Kaolinitic variety	Texturally destructive style; kaolinite, sericite, calcite, pyrite; minor quartz, chlorite, and epidote <i>Summary:</i> Kaolinite + sericite + calcite + pyrite + (quartz) + (chlorite) + (epidote)
Propylitic variety	Chlorite, actinolite, epidote, and calcite replacement of ferromagnesian minerals; oligoclase/albite, epidote, and calcite replacement of plagioclase and K-feldspar; pyrite, specular hematite, minor magnetite <i>Summary:</i> Chlorite + actinolite + epidote + calcite + albite + (pyrite) + (hematite)
Vein style	
Magnetite stringers	± (K-feldspar)
Biotite stringers	Biotite replaced by chlorite
K-feldspar veins	K-feldspar + albite + (calcite) + (epidote) + minor magnetite + minor apatite + (titanite)
Pegmatitic “barren” veins	K-feldspar ± (biotite) replaced by (chlorite)
Pegmatitic bornite-chalcocopyrite veins	K-feldspar ± (biotite) replaced by (chlorite)
Pegmatitic chalcocopyrite-pyrite veins	K-feldspar ± (biotite) replaced by (chlorite)
Bornite-chalcocopyrite veins	None
Chalcocopyrite stringers	Chlorite ± (K-feldspar) + (albite)
Magnetite-sulfide veins	K-feldspar + epidote + (biotite) + (magnetite) ± (chlorite)
Hematite-sulfide veins	K-feldspar + epidote + (biotite) ± (chlorite)
K-feldspar-epidote veins	K-feldspar + epidote + (chlorite)
Chlorite veins	Chlorite + pyrite + chalcocopyrite + (magnetite)
Chlorite-sulfide-calcite veins	K-feldspar
Epidote veins	K-feldspar ±(Albite)
Albite-K-feldspar-scapolite veins	Scapolite + K-feldspar + albite
Sericite veins	Sericite
Quartz veins	None
Zeolite veins	None
Calcite veins	Hematite ± K-feldspar ± albite

Reference: Stanley and others (1995)

Table 4. Alteration-mineral assemblages in the Copper Mountain (British Columbia, Canada) and El Teniente (Chile) porphyry copper deposits.—Continued

Pervasive style	Minerals
El Teniente, Chile	
Vein style	
Magnetite veins	Magnetite + quartz + anhydrite + actinolite + calcic plagioclase + (epidote)
Quartz veins	Quartz + tourmaline + sericite + chlorite + magnetite
Quartz-anhydrite-K-feldspar veins	± Na-K-feldspar
Quartz-chlorite-(anhydrite-biotite-Na-K-feldspar) veins	± Na-K-feldspar
Quartz-(anhydrite-sulfide) ± (K-feldspar-chlorite-biotite) veins	± Na-K-feldspar
Anhydrite-quartz-(sulfide) breccias	± Na-K-feldspar
Quartz-sulfide-(anhydrite) veins	± Biotite + (sericite)
Biotite-tourmaline-quartz-anhydrite-sulfide breccias	Chlorite + (sericite)
Sulfide-chlorite-(anhydrite-quartz) breccias	Biotite
Quartz-anhydrite-sulfide veins	Sericite + chlorite + quartz
Quartz-anhydrite-tourmaline-sulfide-gypsum-carbonate veins	Quartz + sericite + chlorite ± tourmaline
Reference: Cannell and others (2005); Klemm and others (2007)	

