

# Seismicity of the Earth 1900–2012

## Sumatra and Vicinity

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### TECTONIC SUMMARY

The plate boundary southwest of Sumatra, Indonesia, is part of a long tectonic collision zone that extends over 8,000 km from Papua, New Guinea, in the east, to the Himalayan front in the west. The Sumatra-Andaman part of the collision zone forms a subduction-zone megathrust boundary, the Sunda Arc (Sunda-Java trench), which accommodates convergence between the Indo-Australia and Sunda plates. This convergence is responsible for the intense seismicity and volcanism in Sumatra. Plate-boundary related deformation is also not restricted to the subduction zone and overriding plate; the subducting Indo-Australia plate actually comprises two somewhat independent plates (India plate and Australia plate), with small amounts of motion relative to one another, that are joined along a broad, actively deforming region producing seismicity up to several hundred kilometers west of the trench. This deformation is exemplified by the recent April 2012 earthquake sequence, which includes the April 11, 2012 M 8.6, and M 8.2 strike-slip events and their subsequent aftershocks.

The relative motion between the Indo-Australia and Sunda plates is rapid, decreasing from roughly 63 mm/yr near the southern tip of Sumatra (Australia plate relative to Sunda plate) to 44 mm/yr north of Andaman Islands (India plate relative to Sunda) and rotating counterclockwise to the northwest, so that relative motion near Jakarta is nearly trench-normal but becomes nearly trench-parallel near Burma. As a result of the rotation of relative plate motion along the strike of the arc and the interaction of multiple tectonic plates, several interrelated tectonic elements compose the Sumatra-Andaman part of the plate boundary. Most strain accumulation and release occurs along the Sunda megathrust of the main subduction zone, where lithosphere of the subducting Indo-Australia plate is in contact with the overlying Sunda plate down to a depth of 60 km. Strain release associated with deformation within the subducting slab is evidenced by deeper earthquakes that extend to depths of less than 300 km on Sumatra and 150 km or less along the Andaman Islands. The increasingly oblique convergence between these two plates moving northwest along the arc is accommodated by crustal seismicity along a series of transform and normal faults. The Sumatra Fault, a major transform structure that bisects Sumatra, accommodates the northwest-increasing lateral component of relative plate motion. East of the Andaman Islands, back-arc spreading in the Andaman Sea produces a zone of distributed normal and strike-slip faulting. Similar to the Sumatra Fault, the Sagaing Fault near Burma also accommodates the strike-slip component of oblique plate motion.

Paleoseismic studies using coral reefs as a proxy for relative land level changes associated with earthquake displacement suggest that the Sunda arc has repeatedly ruptured during relatively large events in the past, with records extending as far back as the 10th century. In northern Sumatra Island, the southern terminus of the 2004 megathrust earthquake rupture area, a cluster of megathrust earthquakes occurred over a 56-year period between A.D. 1390 and 1455 resulting in uplift substantially greater than that caused by the 2004 event. Studies that look at large sheeted deposits of sand on land interpreted as the transport of debris from a tsunami wave also indicate that this region has gone through significant tsunamis in the past, albeit infrequently.

Prior to 2004, the most recent megathrust earthquakes along the Sunda Arc were in 1797 (M 8.7–8.9), 1833 (M 8.9–9.1), and 1861 (M 8.5). Since 2004, much of the Sunda megathrust between the northern Andaman Islands (India plate) and Enggano Island, a distance of more than 2,000 km, has ruptured in a series of large subduction zone earthquakes—most rupturing the plate boundary south of Banda Aceh. The great M 9.1 earthquake of December 26, 2004, which produced a devastating tsunami, ruptured much of the boundary between Burma and Simeulue Island offshore Banda Aceh. Immediately to the south of the great 2004 earthquake, the M 8.6 Nias Island earthquake of March 28, 2005, ruptured a 400-km section between Simeulue and Batu Islands. Farther south in the Mentawai Islands, two earthquakes on September 12, 2007 of M 8.5 and M 7.9 occurred in the southern portion of the estimated 1797 and 1833 rupture zone, which extends from approximately Enggano Island to the northern portion of Siberut Island. Smaller earthquakes have also been locally important: a M 7.6 rupture within the subducting plate caused considerable damage in Padang in 2009, and a M 7.8 rupture on October 25, 2010, occurred on the shallow portion of the megathrust to the west of the Mentawai Islands, and caused a substantial tsunami on the west coast of those islands.

In addition to the current seismic hazards along this portion of the Sunda arc, this region is also recognized as having one of the highest volcanic activity levels in the world. One of the most dramatic eruptions in human history was the Krakatau eruption on August 27–27, 1883, a volcano just to the southeast of the island of Sumatra, which resulted in over 35,000 casualties.

Subduction and seismicity along the plate boundary adjacent to Java is fundamentally different from that of the Sumatran-Andaman section. Relative motion along the Java arc is trench-normal (approximately 65–70 mm/yr) and does not exhibit the same strain partitioning and back-arc strike-slip faulting that are observed along the Sumatra margin. Neither has the Java subduction zone hosted similar large magnitude megathrust events to those of its neighbor, at least in documented history. Although this region is not as seismically active as the Sumatra region, the Java arc has hosted low-to intermediate-magnitude extensional earthquakes and deep-focus (300–700 km) events and exhibits a similar if not higher volcanic hazard. This part of the arc has also hosted two large, shallow tsunami earthquakes in the recent past which resulted in high tsunami run-ups along the southern Java coast.

### DATA SOURCES

The earthquakes portrayed on the main map (above) are taken from two sources: (1) the Centennial earthquake catalog (Engdahl and Villaseñor, 2002), and annual supplements for the interval 1900–2007, where the magnitude floor is 5.5 globally, and (2) the NEIC PDE Catalog M4.0 and larger from 1973 to July 19, 2012.

1900–2001 Centennial catalog M=5.5–9.5 (Engdahl and Villaseñor, 2002)  
2002–2012 NEIC catalog M=5.5–9.2  
1964–2001 EHD catalog M=4.0–5.4 (Engdahl, written commun.)  
2002–2012 NEIC catalog M=5.5

The nucleation points of great earthquakes (M≥8.3) are designated with a label showing the year of occurrence. Their rupture areas are shown as pale reddish polygons.

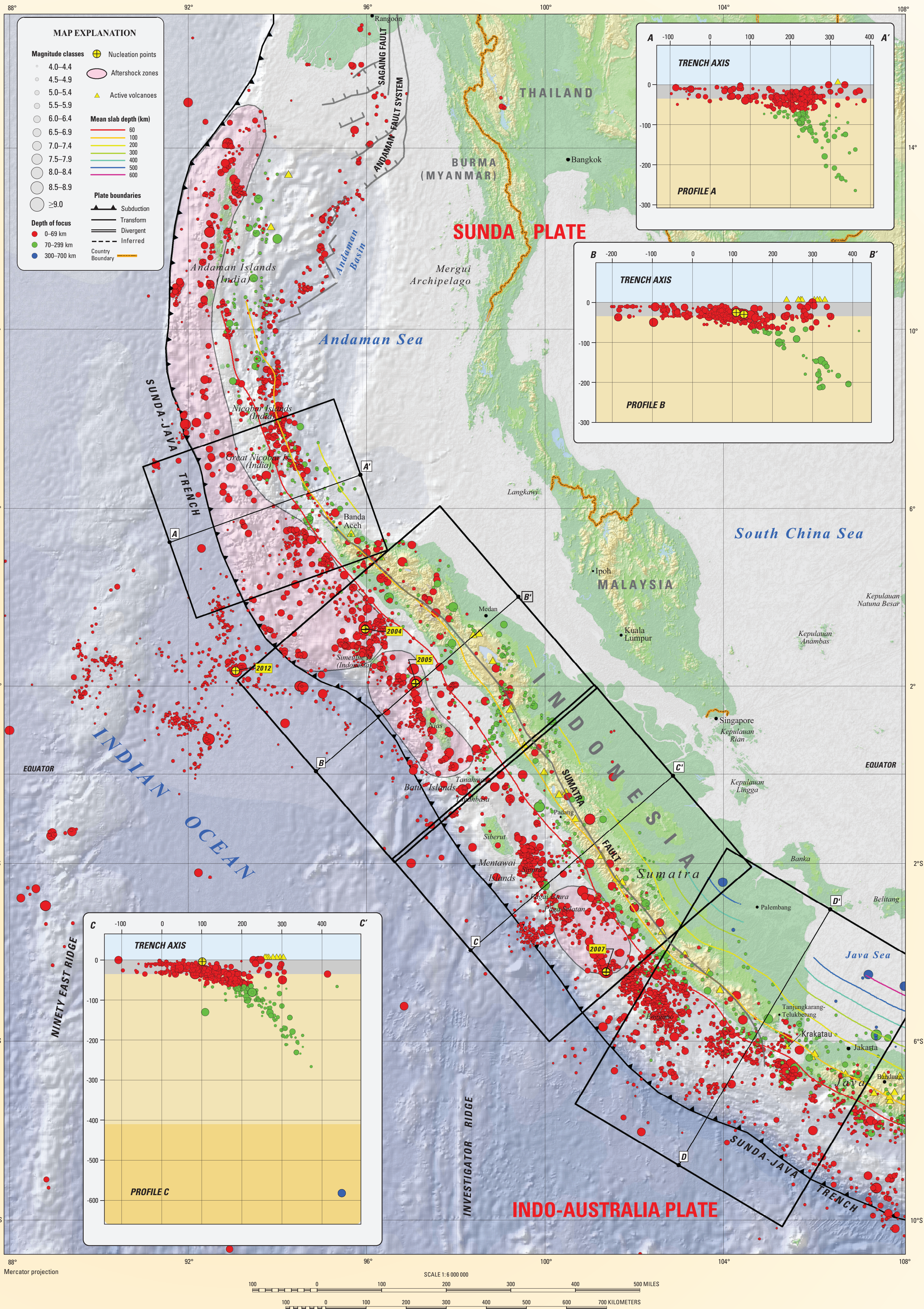
The Seismic Hazard and Relative Plate Motion panel displays the generalized seismic hazard of the region (Giardini and others, 1999) and representative relative plate-motion vectors using the Morvel model (DeMets and others, 2010).

Pre-instrumental seismicity was obtained from the NOAA National Geophysical Data Center (2010) database of significant earthquakes; locations are approximate, based on macro-seismic reports and field investigations. We selected earthquakes with associated reports of moderate to major damage, 10 or more deaths, a estimated magnitude of 7.5 or greater (if known), Modified Mercalli Intensity X, or tsunami generation.

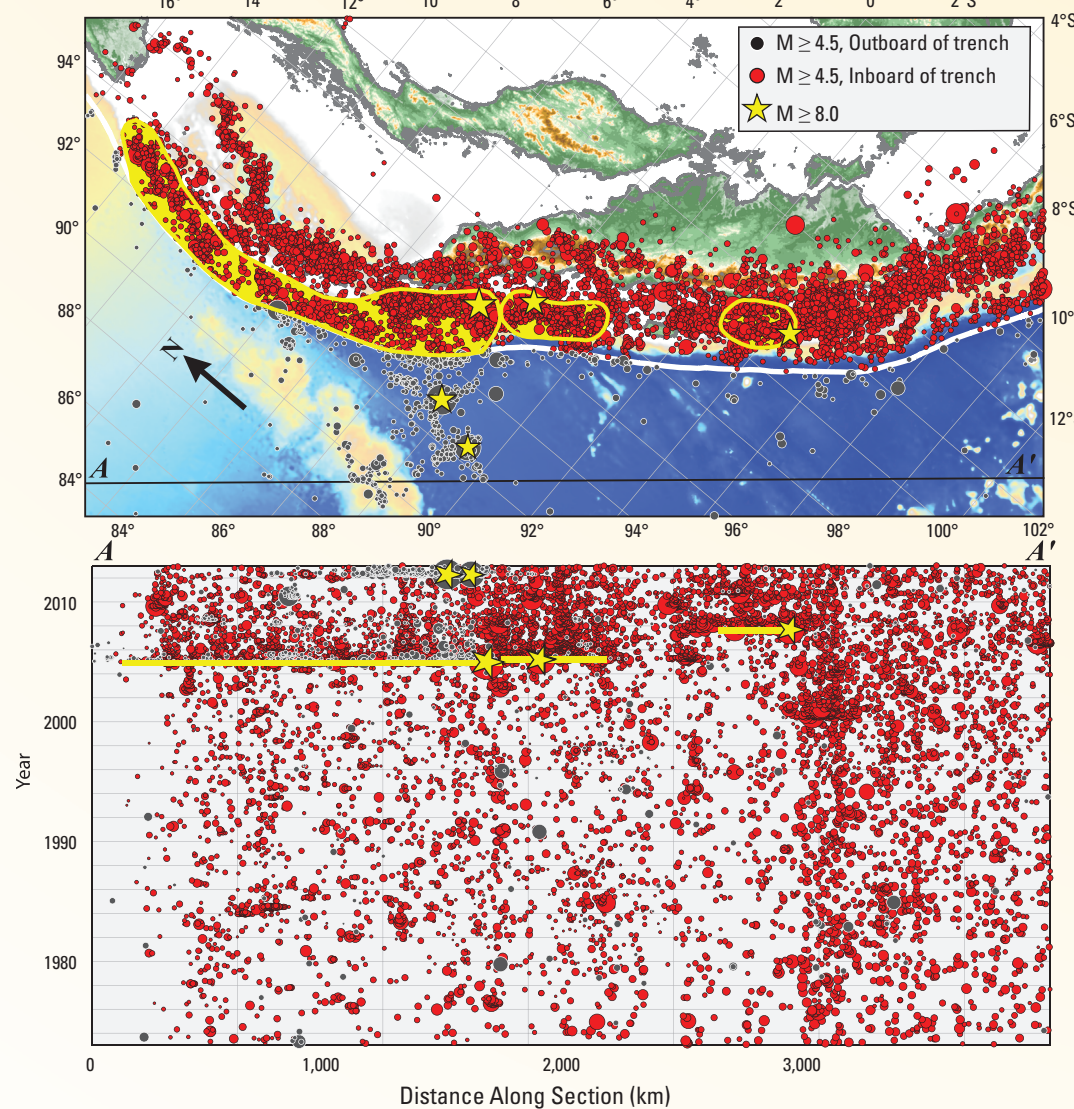
Base map data sources include GEBCO 2008 shaded relief, Volcanoes of the World dataset (Siebert and Simkin, 2002), plate boundaries (Bird, 2003), and geographic information from Digital Chart of the World (1992), and ESR1 (2002). Slab contours are from Hayes and Wald (2010).

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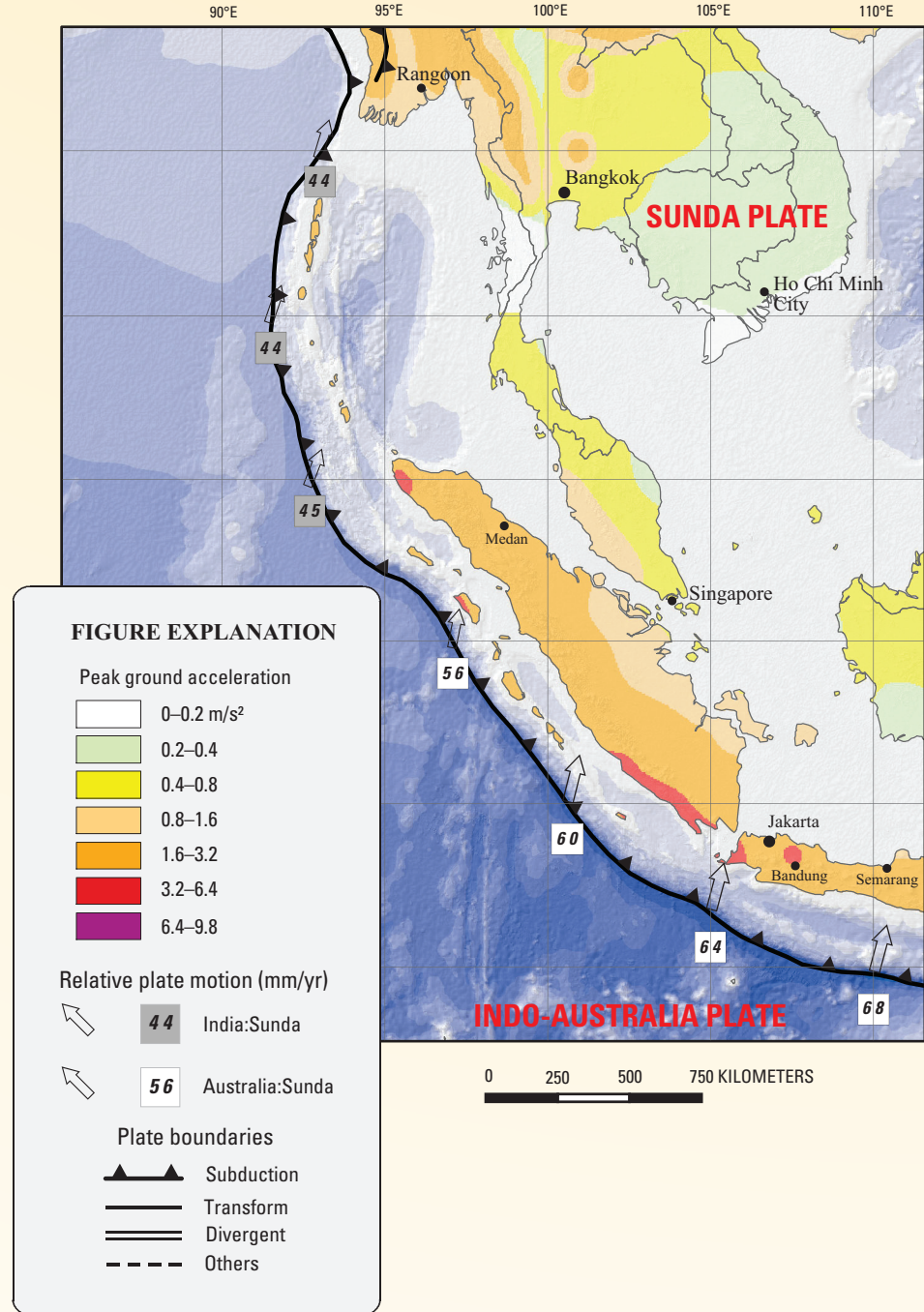


### Time-Space Plot of Seismicity in the Sumatra Arc 1973–2011



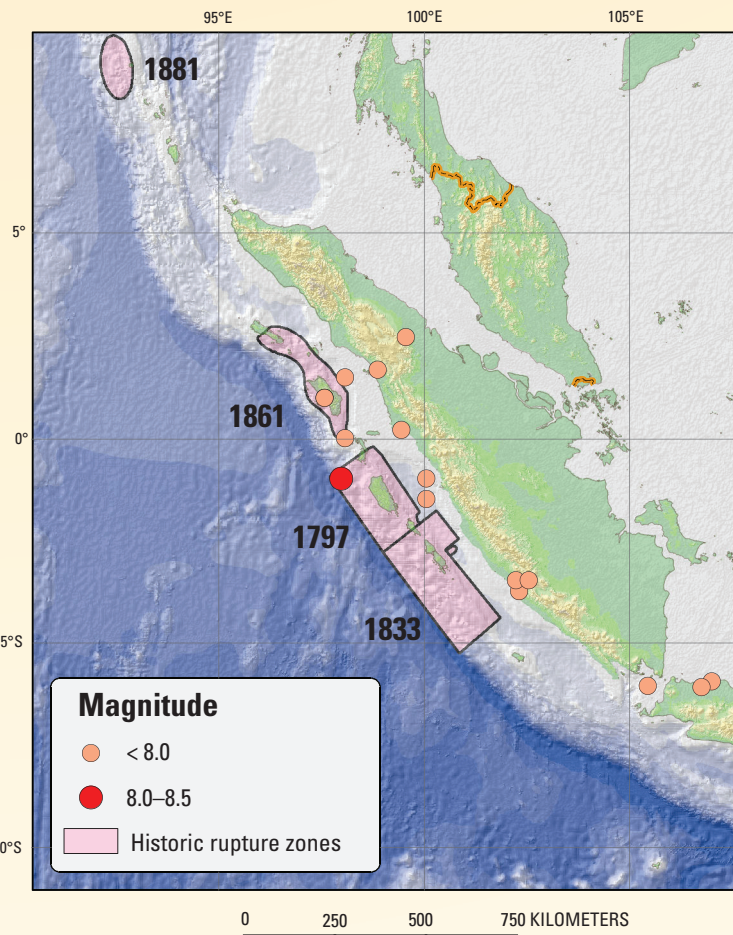
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### Seismic Hazard and Relative Plate Motion



### Pre-Instrumental Seismicity 1500–1899

M ≥ 8.0, tsunami, or deaths



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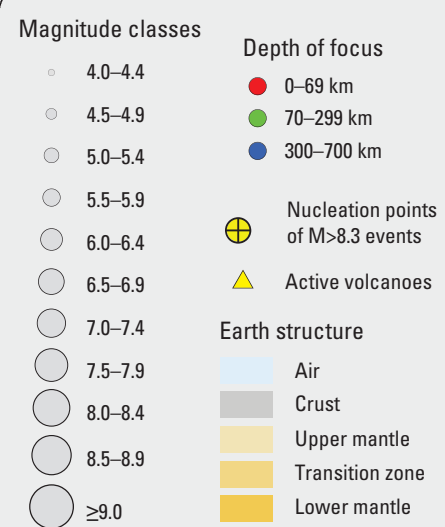
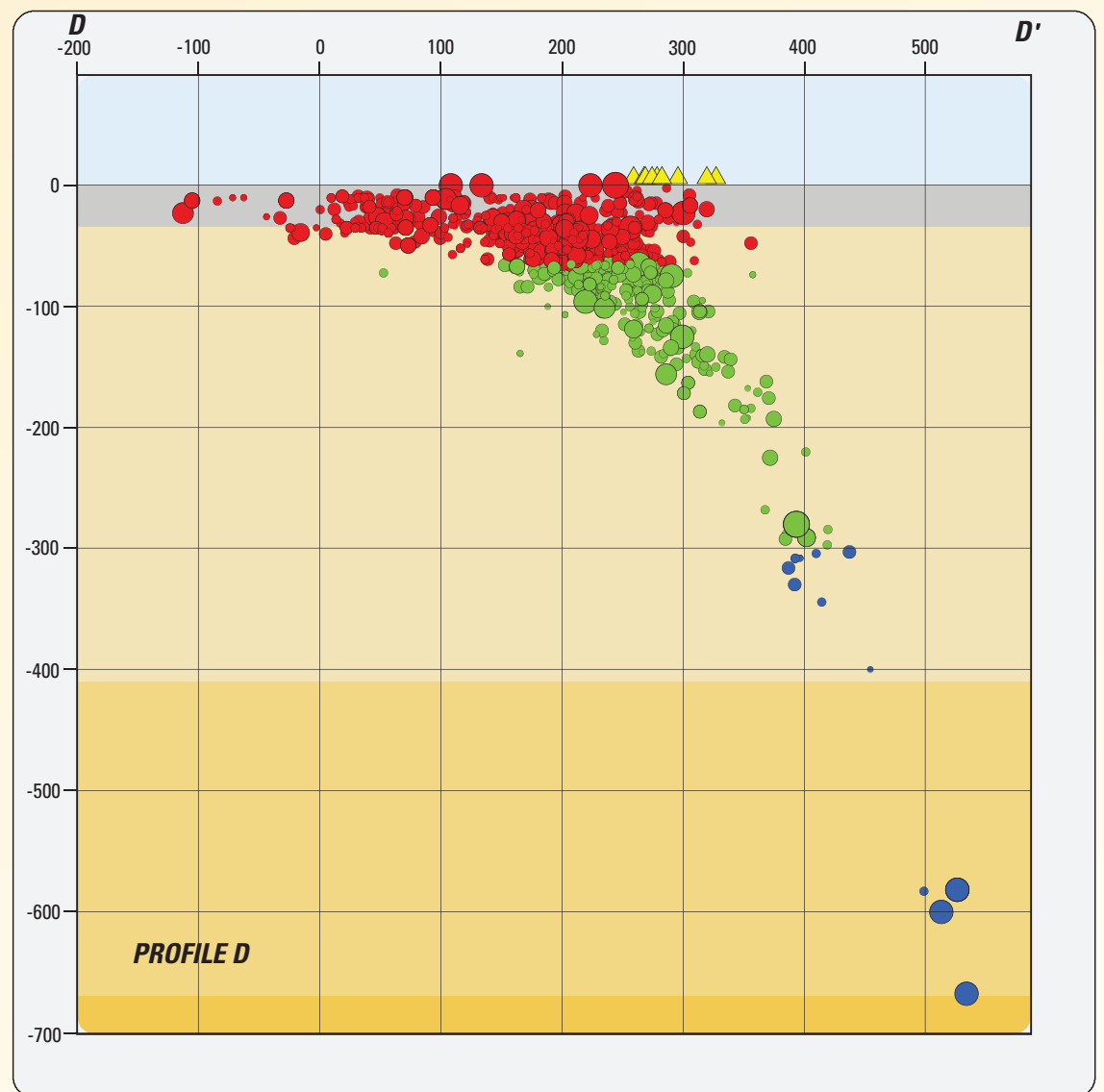
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### DEPTH PROFILE EXPLANATION

Profiles of earthquake and volcano locations are constructed from the mapped data. Locations of the profile intersections with the surface are drawn in the map and labeled to coincide with the profile label. Box defines extent of earthquakes included in the profile. Length of the profile graphic is the same as in the map. Distance in kilometers from the trench axis is indicated in the X direction, depth in kilometers is indicated in the Y direction. There is no vertical exaggeration. See Explanation at side for color key. Not all earthquake depths or magnitudes, are visible in every profile.