



Magnetotelluric Data, Taos Plateau Volcanic Field, New Mexico

By Chad E. Ailes and Brian D. Rodriguez

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Conversion Factors

Inch/Pound to SI

Multiply	By	To obtain
	Length	
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
mile, nautical (nmi)	1.852	kilometer (km)
yard (yd)	0.9144	meter (m)

SI to Inch/Pound

Multiply	By	To obtain
	Length	
meter (m)	3.281	foot (ft)
kilometer (km)	0.6214	mile (mi)
kilometer (km)	0.5400	mile, nautical (nmi)
meter (m)	1.094	yard (yd)

Vertical coordinate information is referenced to the 1866 Clarke Spheroid.

Horizontal coordinate information is referenced to the North American Datum of 1927 (NAD 27).

Elevation, as used in this report, refers to distance above the vertical datum.

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Introduction

The population of the San Luis Basin region of northern New Mexico is growing. Water shortfalls could have serious consequences. Future growth and land management in the region depend on accurate assessment and protection of the region's groundwater resources. An important issue in managing the groundwater resources is a better understanding of the hydrogeology of the Santa Fe Group and the nature of the sedimentary deposits that fill the Rio Grande rift, which contain the principal groundwater aquifers. The shallow unconfined aquifer and the deeper confined Santa Fe Group aquifer in the San Luis Basin are the main sources of municipal water for the region.

The U.S. Geological Survey (USGS) is conducting a series of multidisciplinary studies of the San Luis Basin. Detailed geologic mapping, high-resolution airborne magnetic surveys, gravity surveys, an electromagnetic survey called magnetotellurics (MT), and hydrologic and lithologic data are being used to better understand the aquifers. This report describes a regional east-west MT sounding profile acquired in late July 2009 across the Taos Plateau Volcanic Field where drillhole data are sparse (fig. 1). Resistivity modeling of the MT data can be used to help map changes in electrical resistivity with depths that are related to differences in rock types. These various rock types help control the properties of aquifers. The purpose of this report is to release the MT sounding data collected along the east-west profile. No interpretation of the data is included.

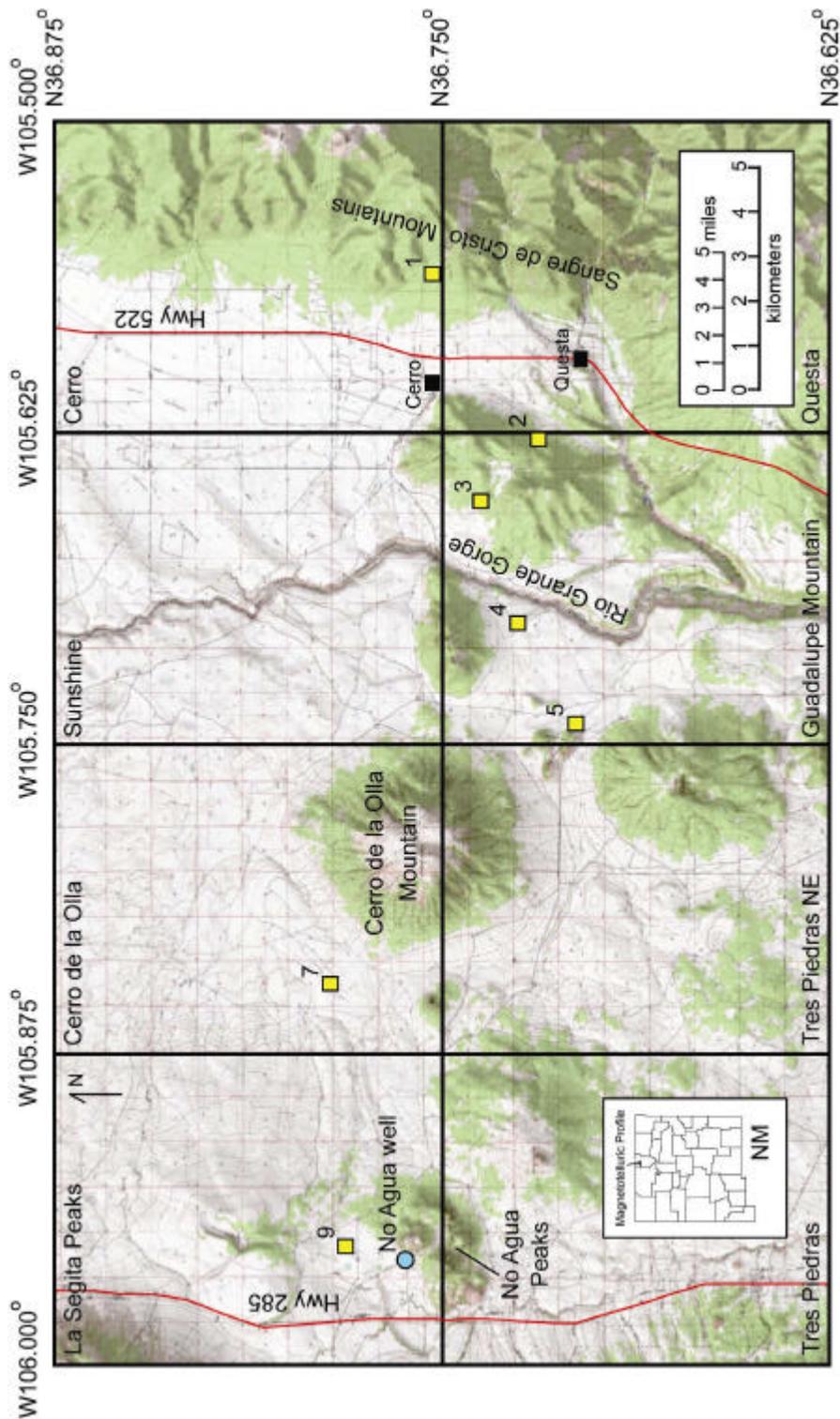


Figure 1. Magnetotelluric profile across the Taos Plateau Volcanic Field, New Mexico. Magnetotelluric stations acquired in late July 2009 are numbered yellow squares (stations 6 and 8 not acquired). Blue circle is No Agua well. Base map from La Segita Peaks, Cerro de la Olla, Sunshine, Cerro, Tres Piedras, Tres Piedras NE, Guadalupe Mountain, and Questa, New Mexico, 1:24,000 topographic quadrangles.

Electrical Rock Properties

Electromagnetic geophysical investigation methods detect variations in the electrical properties of rock units, in particular electrical resistivity, which is measured in units of ohm-meters (Ωm), or its inverse, electrical conductivity in units of Siemens/meter (S/m). Electrical resistivity can be correlated with geologic units on the surface and at depth using lithologic logs to provide a three dimensional picture of subsurface geology. In the upper crust, the resistivities of geologic units are largely dependent upon their fluid content, pore volume porosity, interconnected fracture porosity, and conductive mineral content (Keller, 1987).

Although there is not a one-to-one relationship between lithology and resistivity, there are general correlations that can be made using typical values, even though values can be found at other geographic locations (Palacky, 1987) that may fall outside of the ranges presented below. Fluids within the pore spaces and fracture openings, especially if saline, can reduce resistivities in what would otherwise be a resistive rock matrix (Hallenburg, 1998; Hearst and Nelson, 1985; Hearst and others, 2000; Keller, 1987; Keller and Frischknecht, 1966; Palacky, 1987). Resistivity can also be lowered by the presence of electrically conductive clay minerals, graphitic carbon, and metallic mineralization. It is common, for example, for altered volcanic rocks to contain replacement minerals that have resistivities ten times lower than those of the surrounding rocks (Nelson and Anderson, 1992). Fine grained sediments, such as clay rich alluvium, marine shales, and other mudstones are normally conductive, with resistivities ranging from a few Ωm to tens of Ωm (Keller, 1987; Palacky, 1987). Metamorphic rocks (non graphitic) and unaltered, unfractured igneous rocks are normally moderately to highly resistive (a few hundred to thousands of Ωm). Carbonate rocks can have similarly high resistivities depending on their fluid content, porosity, and impurities (Keller, 1987; Palacky, 1987). Fault zones may be moderately conductive (tens of Ωm) when composed of rocks fractured enough to have hosted fluid transport and consequent mineralogical alteration (Eberhart Phillips and others, 1995). At greater depths, higher subsurface temperatures cause higher ionic mobility that reduces rock resistivities (Keller, 1987; Palacky, 1987). Tables of electrical resistivity for a variety of rocks, minerals, and geological environments may be found in Keller (1989) and Palacky (1987).

Magnetotelluric Method

The magnetotelluric method is a passive surface geophysical technique that uses the Earth's natural electromagnetic fields to investigate the electrical resistivity structure of the subsurface from depths of tens of meters to tens of kilometers (Vozoff, 1991). Natural variations of the Earth's magnetic and electric fields are measured and recorded at each magnetotelluric station. Worldwide lightning activity at frequencies of about 1 to 20,000 Hertz and geomagnetic micro-pulsations at frequencies of about 0.0001 to 1 Hertz provide the majority of the signal sensed by the magnetotelluric method. The natural electromagnetic waves propagate vertically in the Earth because the very large contrast in the resistivity of the air and the Earth causes a vertical refraction of the electromagnetic wave at the surface (Vozoff, 1972).

The orthogonal horizontal electric and magnetic field components (E_x , E_y , H_x , and H_y) and the vertical magnetic field component (H_z) are recorded. The resulting time-series signals are used to derive tensor apparent resistivities and phases after first converting them to complex cross-spectra using fast Fourier transform (FFT) techniques and least-squares, cross-spectral analysis (Bendat and Piersol, 1971) to solve for a tensor transfer function. If one assumes that the Earth consists of a two-input, two-output linear system in which the orthogonal magnetic fields are input and the orthogonal electric fields are output, then a transfer function can be calculated that relates the observed electric fields to the magnetic fields. For resistivity modeling, magnetotelluric data are normally rotated into directions that

are parallel and perpendicular to the subsurface geologic strike. Subsurface geologic strike can be estimated by determining the horizontal direction (H_{xy} or H_{yx}) that the vertical magnetic field (H_z) “tips” (the Tipper strike direction).

For a two-dimensional Earth, the MT fields can be decoupled into transverse electric (TE) and transverse magnetic (TM) modes. Two-dimensional resistivity modeling generally is computed to fit both modes. When the geology satisfies the two-dimensional assumption, the MT data for the TE mode are assumed to represent the electric field oriented along geologic strike, and the data for the TM mode are assumed to represent the electric field oriented across strike. The MT method is well suited for studying complicated geological environments because the electric and magnetic relations are sensitive to vertical and horizontal variations in resistivity. High-resolution shallow subsurface characterization is possible for closely spaced MT stations, but the resolution of the subsurface decreases for deeper measurements and for widely-spaced stations. The method is capable of establishing whether the electromagnetic fields are responding to subsurface rock bodies of effectively 1, 2, or 3 dimensions. An introduction to the magnetotelluric method and references for a more advanced understanding are in Kaufman and Keller (1981), Dobrin and Savit (1988), and Vozoff (1991).

Magnetotelluric Survey

Seven MT soundings were collected in July 2009 along a 35-km-long east-west profile in northern New Mexico (fig. 1). The profile starts north of the No Agua Peaks about 2.5 kilometer east of State Highway 285, then continues easterly, until a bend south of Cerro de la Olla mountain, then continues east across the Rio Grande Gorge, and ends in the Sangre de Cristo Mountains about 3 km east of State Highway 522, east of the town of Cerro and northeast of the town of Questa, New Mexico. The profile location was selected to map electrical resistivity with depth related to the properties of aquifers beneath the Taos Plateau Volcanic Field.

Station locations were chosen for proximity to roads and in order to avoid electrical noise from power lines. All data at the stations were collected with a portable Electromagnetic Instruments, Inc. (EMI), MT-1 system (Electromagnetic Instruments, Inc., 1996). Horizontal electric fields were recorded using copper sulfate electrodes placed in an L-shaped, three-electrode array with dipole lengths of 30 meters (m). The orthogonal horizontal magnetic fields were measured in the direction of the electric-field array and were sensed using high magnetic permeability mu-metal-cored induction coils. Frequencies were sampled from about .01 Hertz to about 100 Hertz at each station. Station 9 was located close to the No Agua well to help calibrate future resistivity modeling of the MT data. The No Agua well is a 354 meters test well with lithologic descriptions that encountered granite at 350 meters feet. Table 1 lists the seven MT station locations. Stations 6 and 8 were not acquired because of site access problems. Remote reference stations were not used because manmade noise sources were assumed to be minimal in this remote area.

Table 1. Magnetotelluric station coordinates.

[Coordinates are referenced to the 1866 Clarke spheroid and North American 1927 Western United States datum. Longitude and latitude format is degrees:minutes:seconds. Universal Transverse Mercator (UTM) units and station elevations are in meters. The accuracy of the north and east component is ± 5 m and is ± 10 m for the elevation. X direction (X Dir.) is in degrees clockwise from true north]

Station	X Dir.	Longitude	Latitude	North (m)	East (m)	Elevation (m)
01	295	-105:33:40	36:45:11	4,067,415	449,925	2,608
02	004	-105:37:38	36:43:09	4,063,693	443,996	2,325
03	171	-105:39:08	36:44:15	4,065,746	441,778	2,353
04	041	-105:42:05	36:43:32	4,064,453	437,367	2,300
05	005	-105:44:30	36:42:25	4,062,421	433,753	2,317
07	279	-105:50:48	36:47:10	4,071,274	424,446	2,363
09	141	-105:57:08	36:46:52	4,070,799	415,019	2,537

Magnetotelluric Data

The recorded time-series data were converted to the frequency domain and processed to determine the impedance tensor, which is used to derive apparent resistivities and phases at each site. Rotation of the impedance tensor allows for decoupling into the TE and TM modes. The data provided here have not been rotated from the original acquisition orientation (X Dir.) listed in table 1 above. During the analysis and interpretation process, each station should be rotated to a fixed angle determined by the given nominal profile orientation. Cross-power files were sorted to select optimal signal-to-noise time-series data sets (see appendix 1).

Cultural features, such as fences, pipelines, communication lines, moving vehicles and trains, and other manmade sources of electromagnetic noise can contaminate the responses of the MT system. Care is taken to avoid these sources of noise when acquiring the data.

The figures in appendix 1 represent the field-processed magnetotelluric data for each station, after the time-series data were converted to the frequency domain and the tensor-transfer function was developed. They include some data scatter and poor signal-to-noise ratios. The only effort aimed at removing noisy data points was to visually inspect and select the best signal-to-noise field data to combine into the final data plots.

For each station, eight separate plots are given:

1. Apparent Resistivity (x and o symbols are xy and yx components)
2. Impedance Phase (x and o symbols are xy and yx components)
3. Impedance Skew
4. Multiple Coherency (x and o symbols are xy and yx components)
5. Impedance Polar Plots
6. Tipper Magnitude
7. Tipper Strike
8. HzHx (x symbol) and HzHy (o symbol) Coherency

Error bars (|,|) on the Apparent Resistivity, Impedance Phase, Impedance Skew, Tipper Magnitude, and Tipper Strike plots represent probable errors within one standard deviation of the sample variance (Gamble and others, 1979).

Apparent resistivity is the ratio at a given frequency of the electric field strength magnitude to the magnetic field strength magnitude. The impedance phase is proportional to the slope of the apparent resistivity curve on a log-log plot, relative to a baseline at -45° (Vozoff, 1991). A measure of the dimensionality for magnetotelluric data is provided by the impedance skew of the impedance tensor (Vozoff, 1972). If the effective, measured resistivity response to the geology beneath an magnetotelluric station is truly one- or two-dimensional, then the skew will be zero. Both instrumental and environmental sources of noise contribute to non zero skew values but are typically small (about 0.1) for relatively low-noise-level recordings. Higher skews (more than 0.2) indicate either the resistivity response to 3-D geology or higher levels of noise.

In the study area, noise from a number of small powerlines and small moving vehicles was negligible at distances of 0.25 km and farther from the noise source. Powerline amplitude levels were measured at each site and were typically less than 20 percent of the maximum recordable signals. Noise from larger power lines, power generators, pipelines, and trains was negligible at distances greater than 5 km. Local lightning, wind, and rainstorms also can degrade data quality, but these noise sources were avoided by not recording during active thunderstorm periods. Burying the magnetic induction coils and keeping the electric dipole wires flat on the ground helped to minimize wind noise.

Predicted values of the electric field can be computed from the measured values of the magnetic field (Vozoff, 1991). The coherence of the predicted electric field with the measured electric field is a measure of the signal-to-noise ratio provided in the multiple coherency plots. Values are normalized between 0 and 1, where values at 0.5 signify signal levels equal to noise levels. For this data set, coherencies were generally above 0.75, except in the middle frequencies (from about 0.1 Hertz to 5 Hertz) at most stations, and also in the lower frequencies for station 7, where a nearby thunderstorm made it difficult to collect low-frequency data with high amplifier gains.

The figures in appendix 1 represent the field-processed magnetotelluric data at each station and include some data scatter and poor signal-to-noise ratios. The only effort aimed at removing noisy data points was to visually inspect and digitally select the best signal to noise field data to combine into the final data plots.

The impedance polar plots provide a measure of the magnetotelluric data dimensionality (Reddy and others, 1977). For 1-D resistivity structures, the principal impedance (off diagonal elements) polar diagram (dashed line) is a circle. For 2-D or 3-D resistivity structures, the principal impedance polar diagram (dashed line) elongates either parallel to or perpendicular to strike direction. Over resistors, the principal impedance polar diagram elongates perpendicular-to-strike direction, while over conductors, the principal impedance polar diagram elongates parallel-to-strike direction. For 2-D resistivity structures, the additional impedance polar diagram (solid line) attains the shape of a symmetric clover leaf. For 3-D resistivity structures, the additional impedance polar diagram (solid line) elongates in one direction, and its amplitude is comparable to that of the principal impedance polar diagram (dashed line), although high noise levels can produce the same effect on the polar diagram. A 3-D analysis of polar plots at each frequency should also take into account the corresponding coherence and skew values along with their associated error levels. The polar plots computed for the data show the electromagnetic response for all stations was 3-D over the middle to low frequencies measured at acceptable noise levels. The electromagnetic response for stations 2, 3, and 4 was 3-D over all frequencies.

The tipper can be calculated from the vertical component of the magnetic field. The tipper magnitude is a measure of the tipping of the magnetic field out of the horizontal plane (Vozoff, 1991). The magnitude is 0 for the 1-D case, typically increases between 0.1 to 0.5, and rarely is as large as 1 as it responds to vertical and subvertical structures. The tipper strike typically is used to help resolve the 90° ambiguity in the impedance rotation angle. The tipper magnitude of these stations were all above 0.1 indicating vertical and sub-vertical structure at depth.

The HzHx and HzHy coherency is a measure of the signal-to-noise ratio of the vertical magnetic field with respect to each of the orthogonal, horizontal magnetic field directions. Values are normalized between 0 and 1, where values of 0.5 signify signal levels equal to noise levels. These three components of magnetic-field coherence provide a check on the quality of the measured values in the tipper magnitude and tipper strike plots.

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Appendix 1. Magnetotelluric Data Plots

There are eight separate plots for each station:

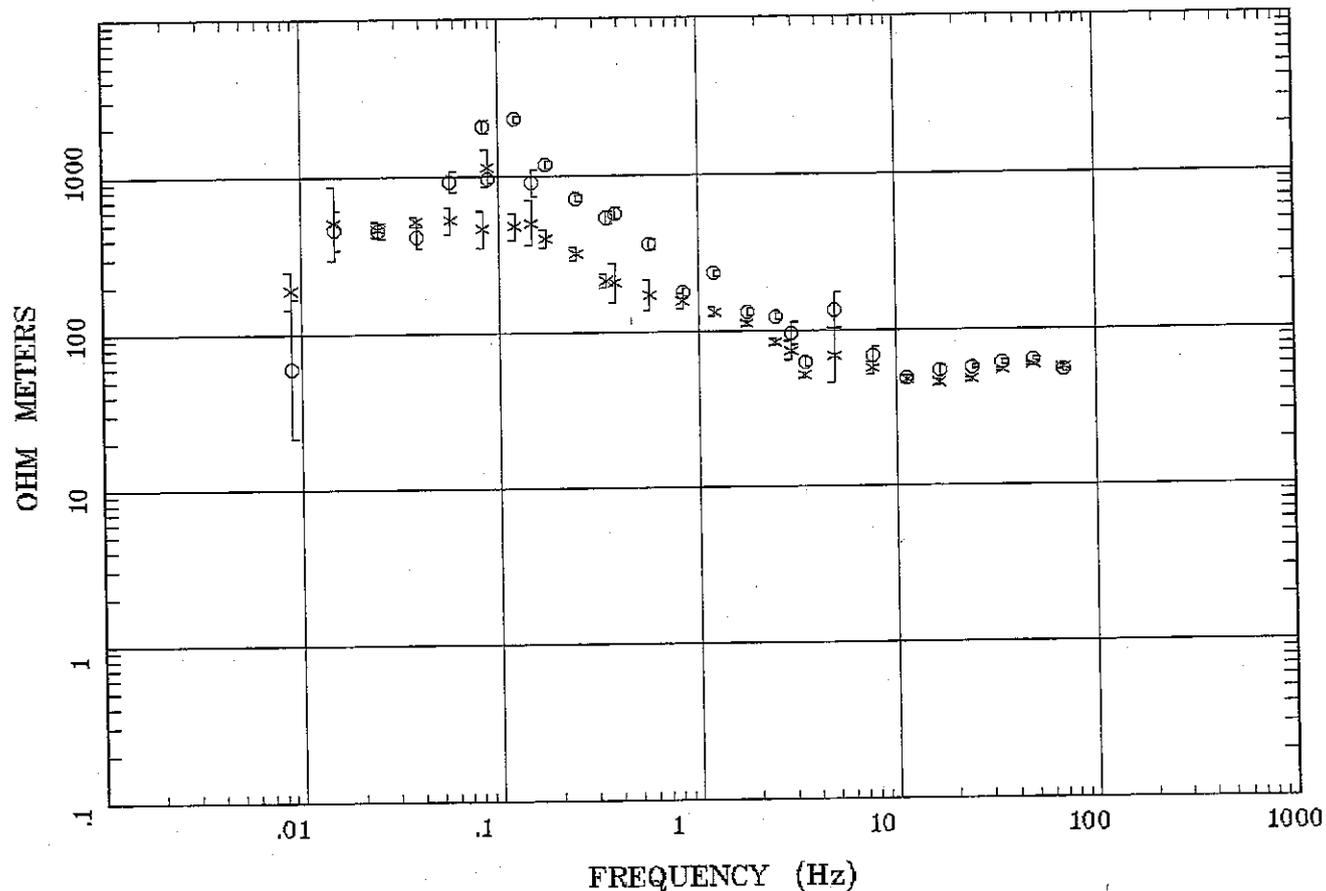
1. Apparent Resistivity for the unrotated xy (x symbol) and yx (o symbol) modes
2. Impedance Phase for the unrotated xy (x symbol) and yx (o symbol) modes
3. Impedance Skew for the impedance tensor
4. Multiple Coherency for the xy (x symbol) and minimum (o symbol) modes of the electric field
5. Impedance Polar Plots (at 12 selected frequencies)
6. Tipper Magnitude for the vertical magnetic field
7. Tipper Strike for the vertical magnetic field
8. HzHx (x symbol) and HzHy (o symbol) Coherency

Error bars (|,|) on the apparent resistivity, impedance phase, skew, tipper magnitude, and tipper strike plots represent probable errors within one standard deviation of the sample variance (Gamble and others, 1979). Refer to the “Magnetotelluric Data” section in this report for an explanation of these plots.

APPARENT RESISTIVITY

Wheeler Peak Quad, NM

Station 1



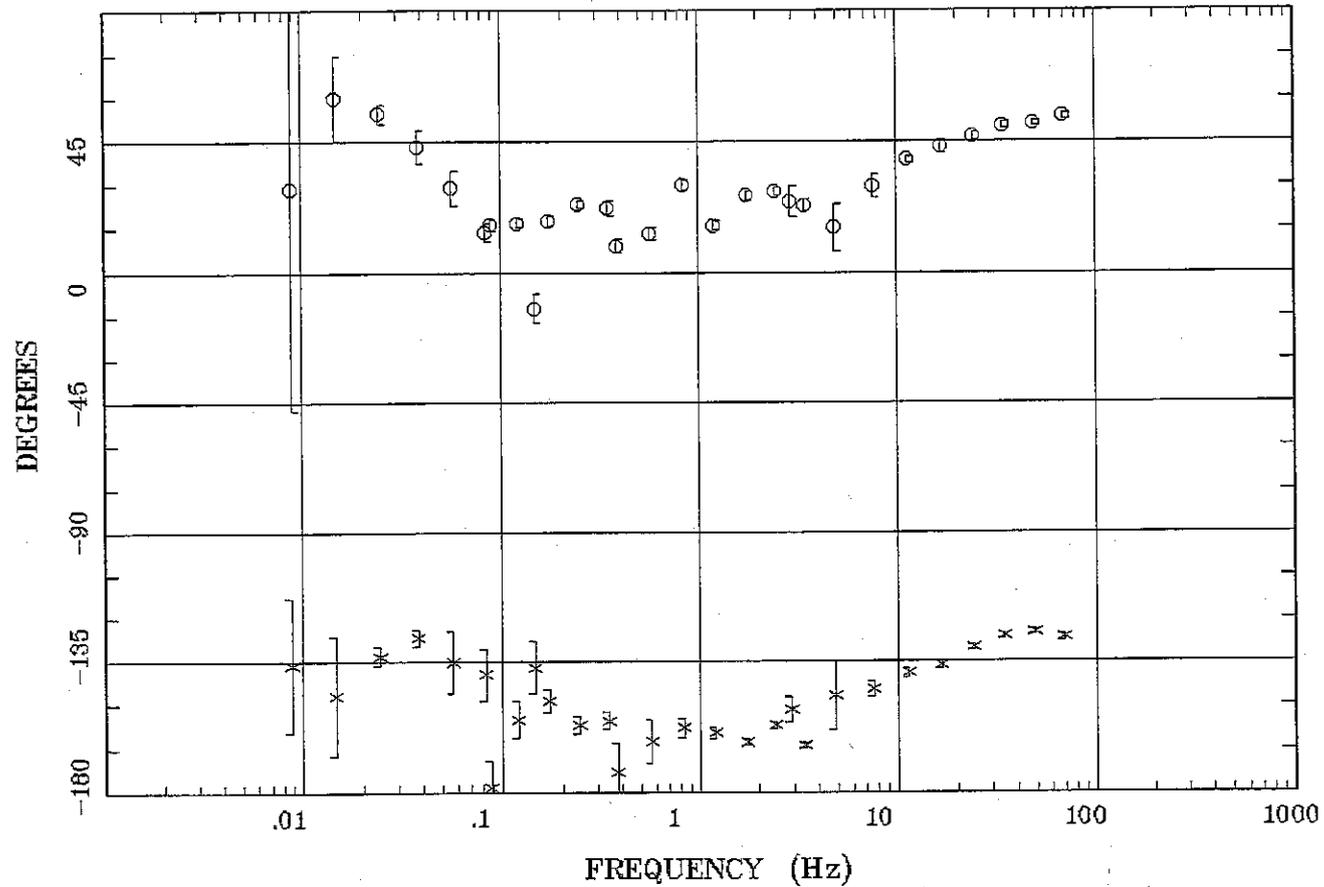
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 Survey Co:USGS

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Station 1

IMPEDANCE PHASE

Wheeler Peak Quad, NM



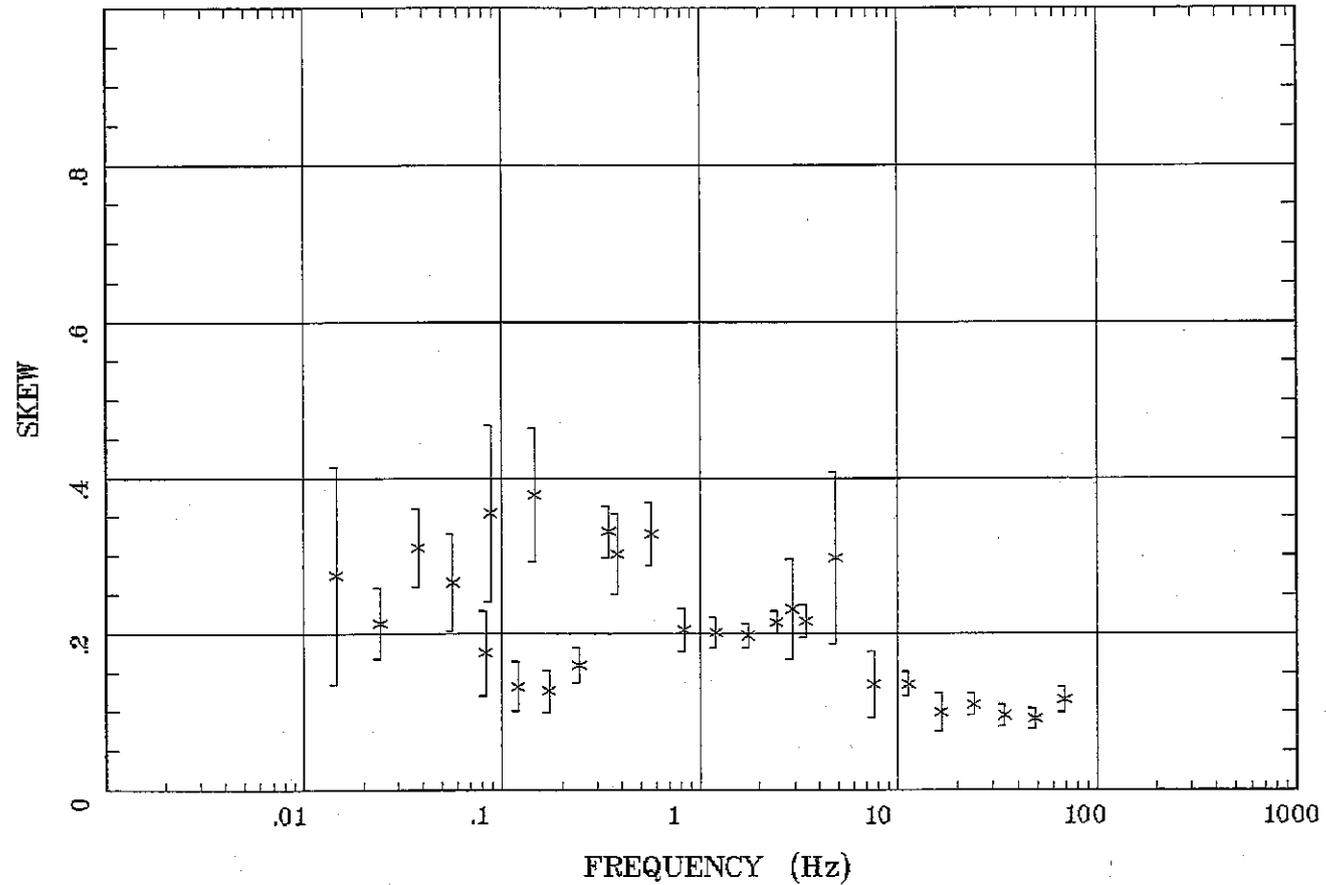
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Station 1

IMPEDANCE SKEW

Wheeler Peak Quad, NM



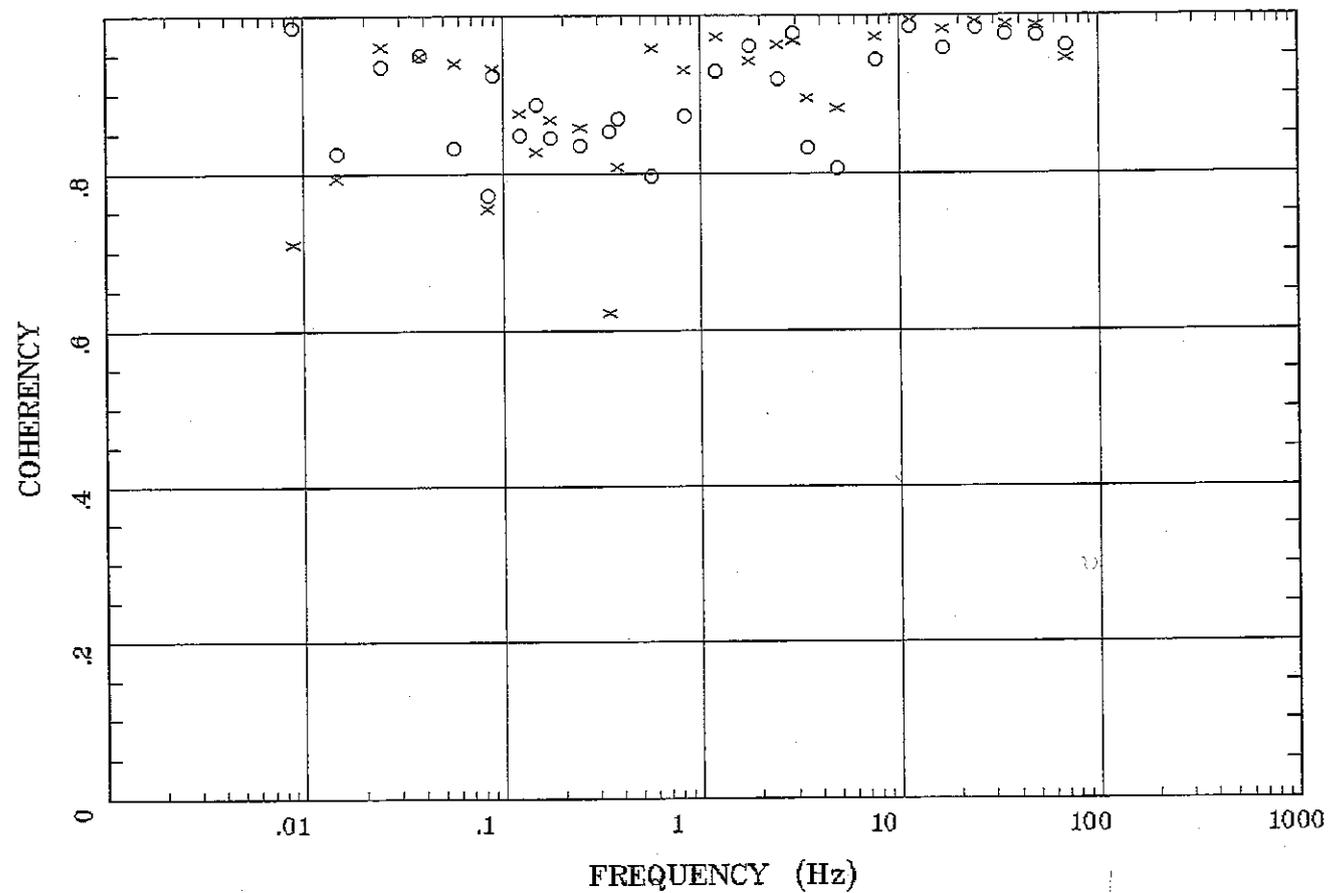
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Station 1

E MULT Coh.

Wheeler Peak Quad, NM



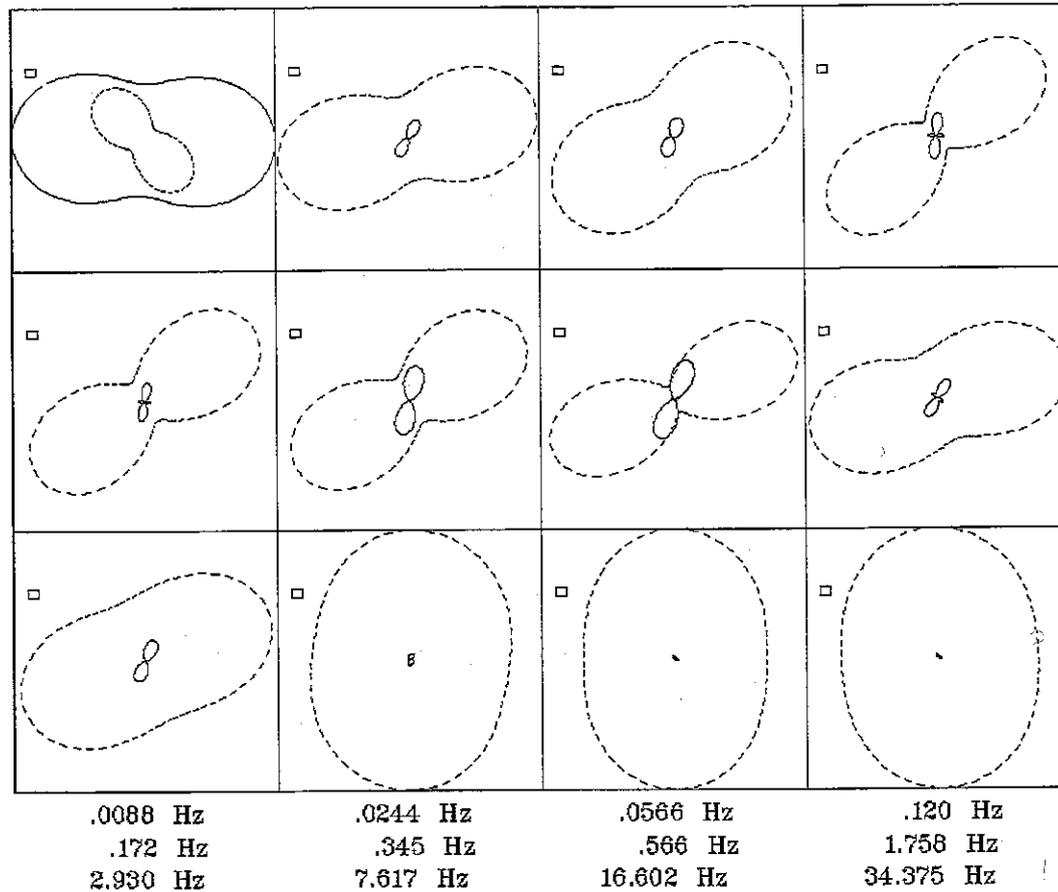
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Station 1

POLAR PLOTS

Wheeler Peak Quad, NM



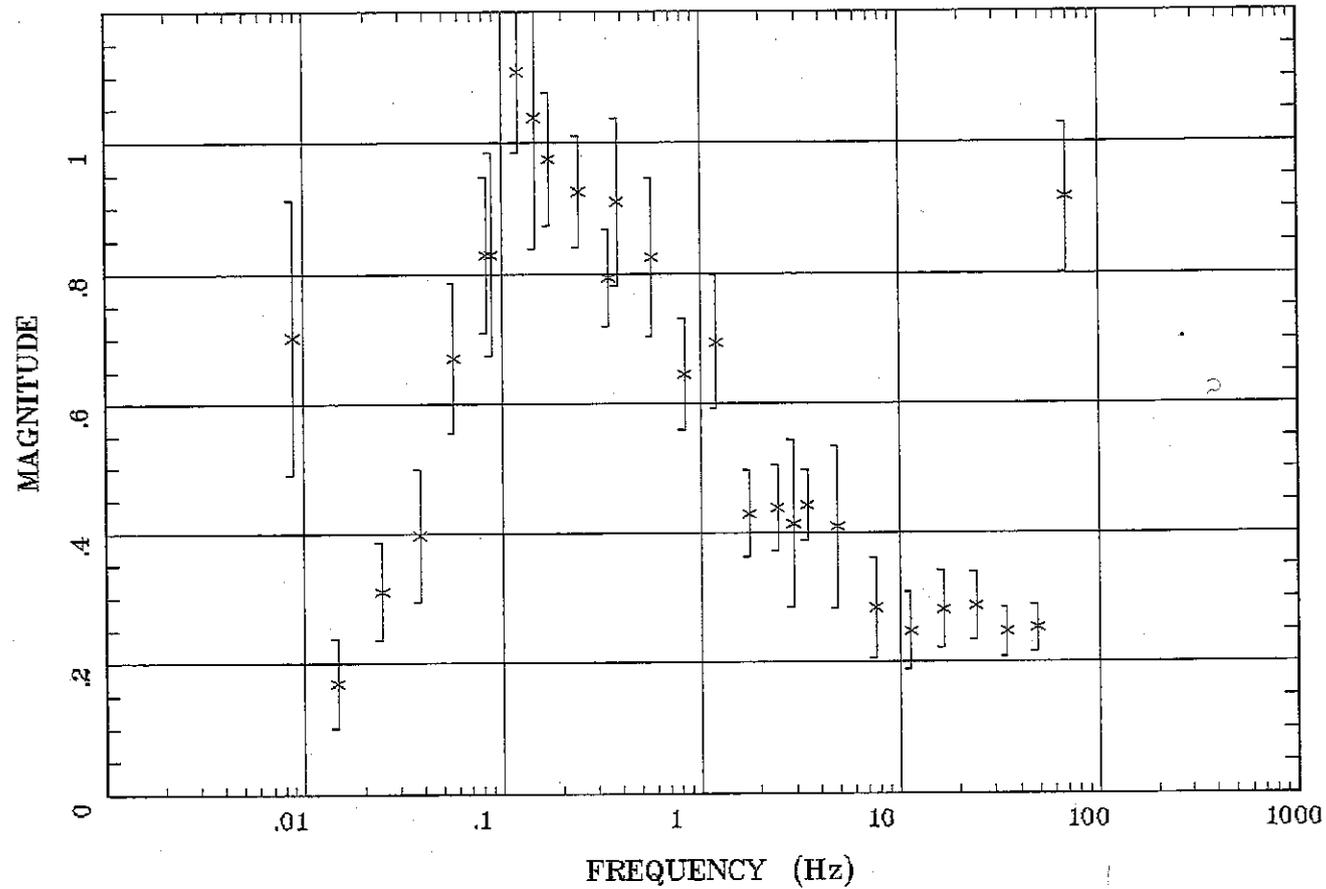
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Station 1

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



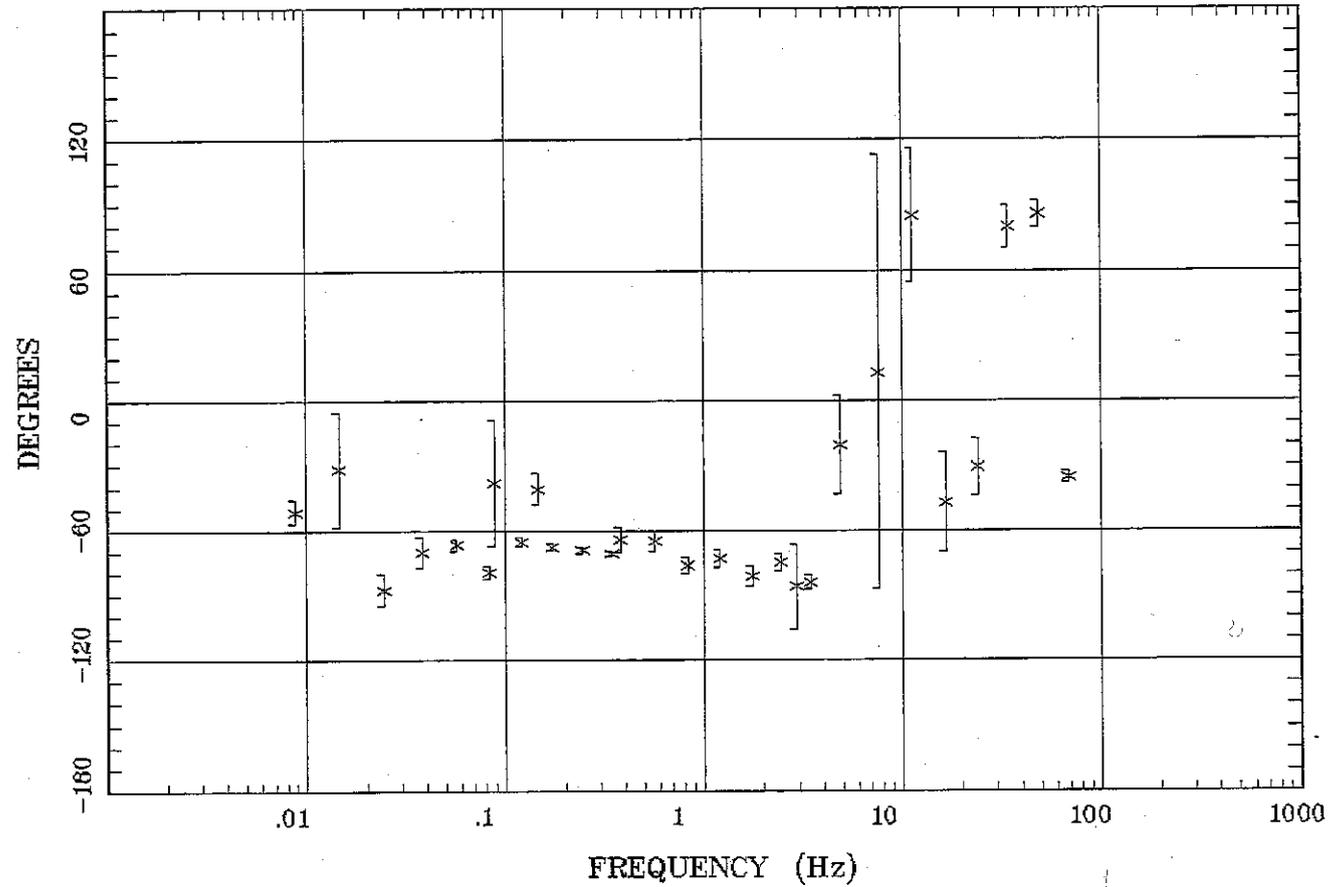
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Station 1

TIPPER STRIKE

Wheeler Peak Quad, NM



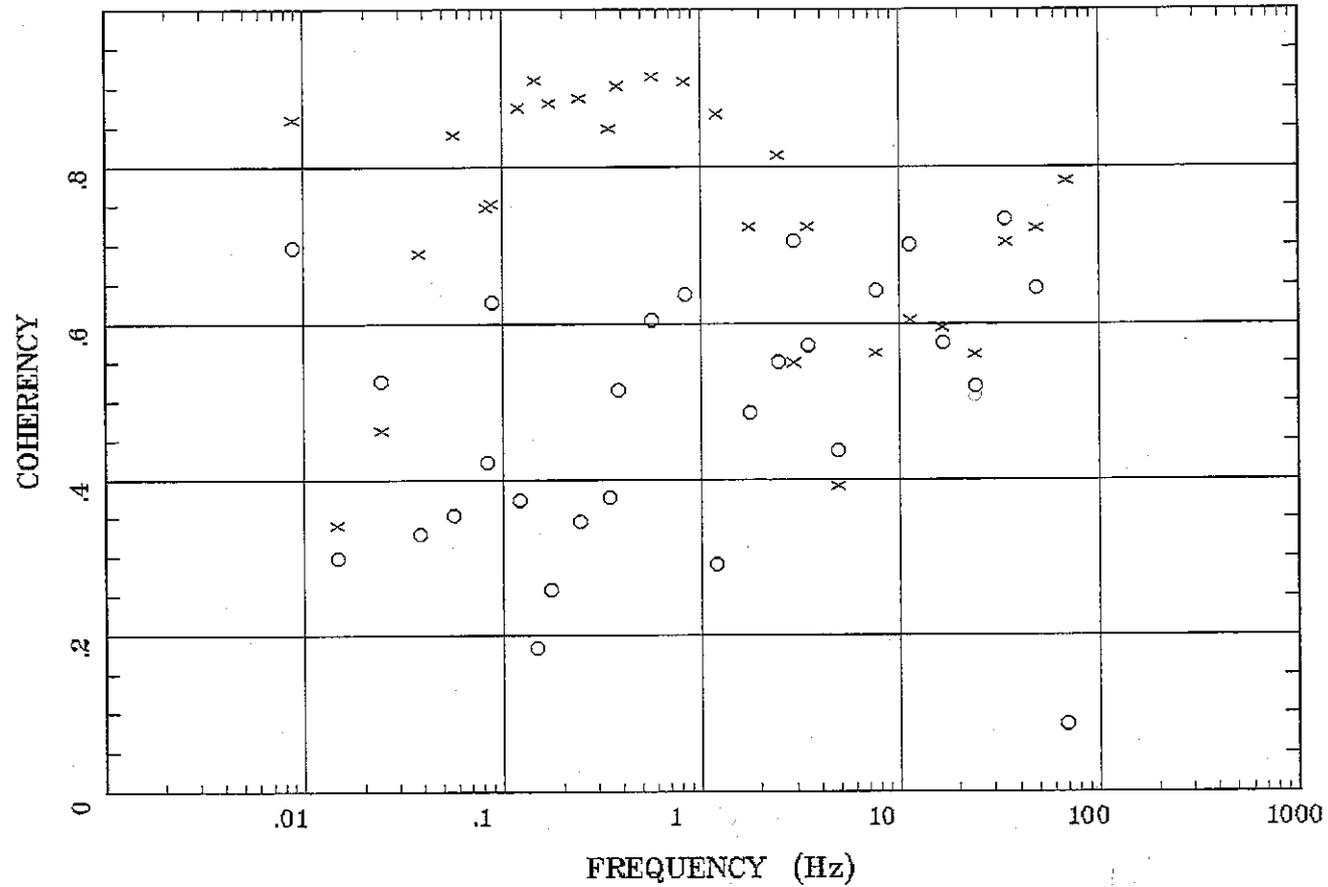
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Station 1

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



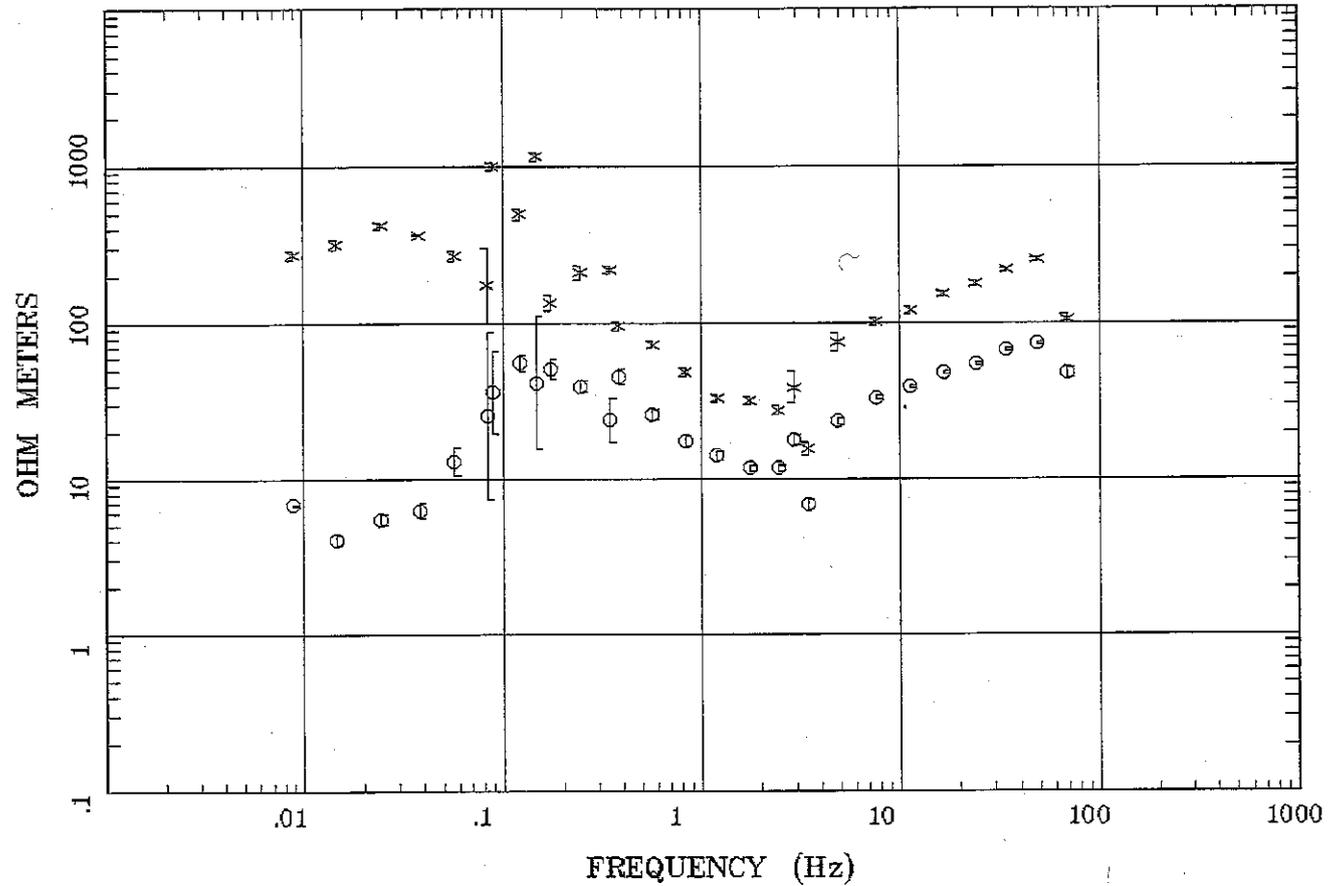
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Station 2

APPARENT RESISTIVITY

Wheeler Peak Quad, NM



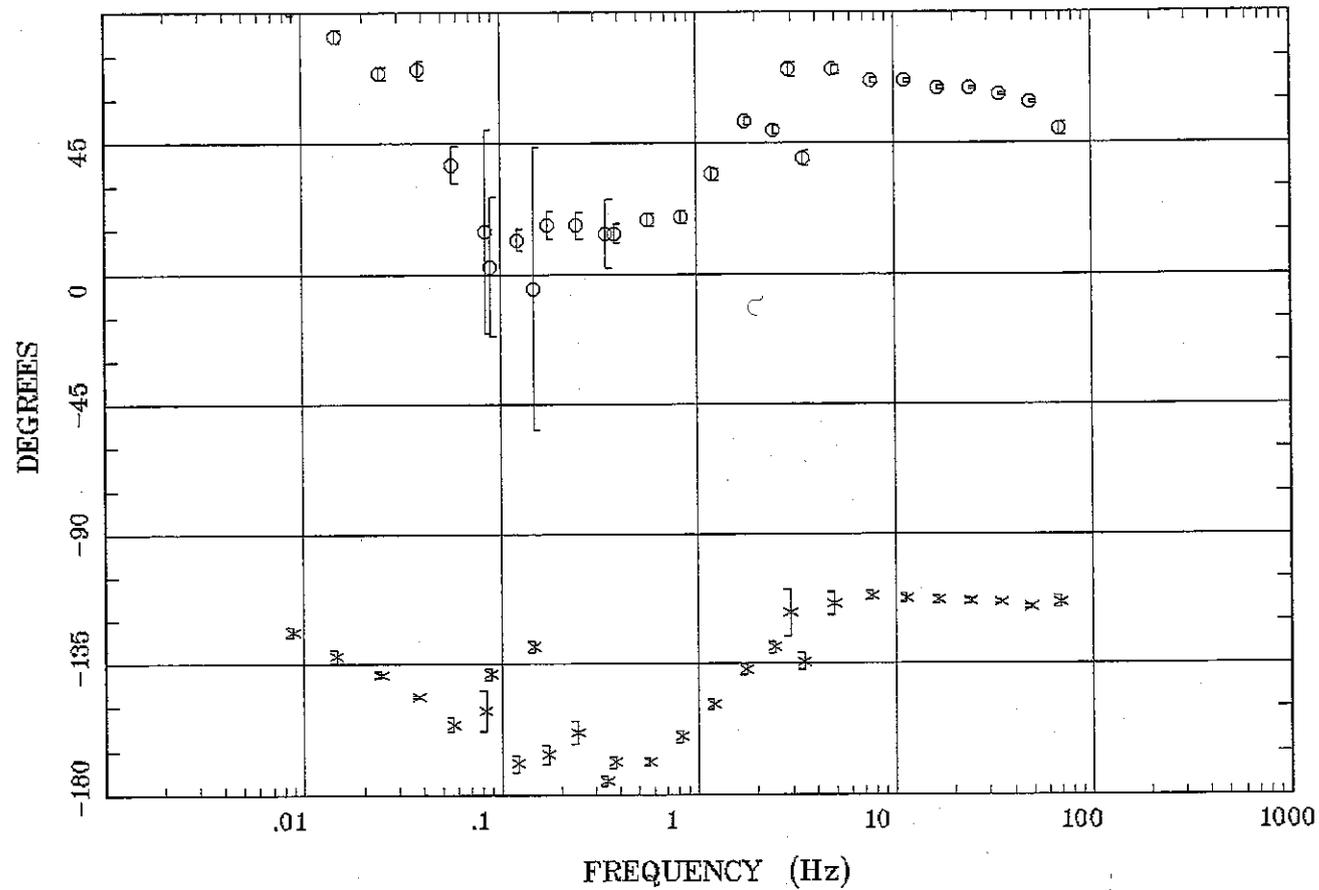
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< EMI - ElectroMagnetic Instruments >

Station 2

IMPEDANCE PHASE

Wheeler Peak Quad, NM



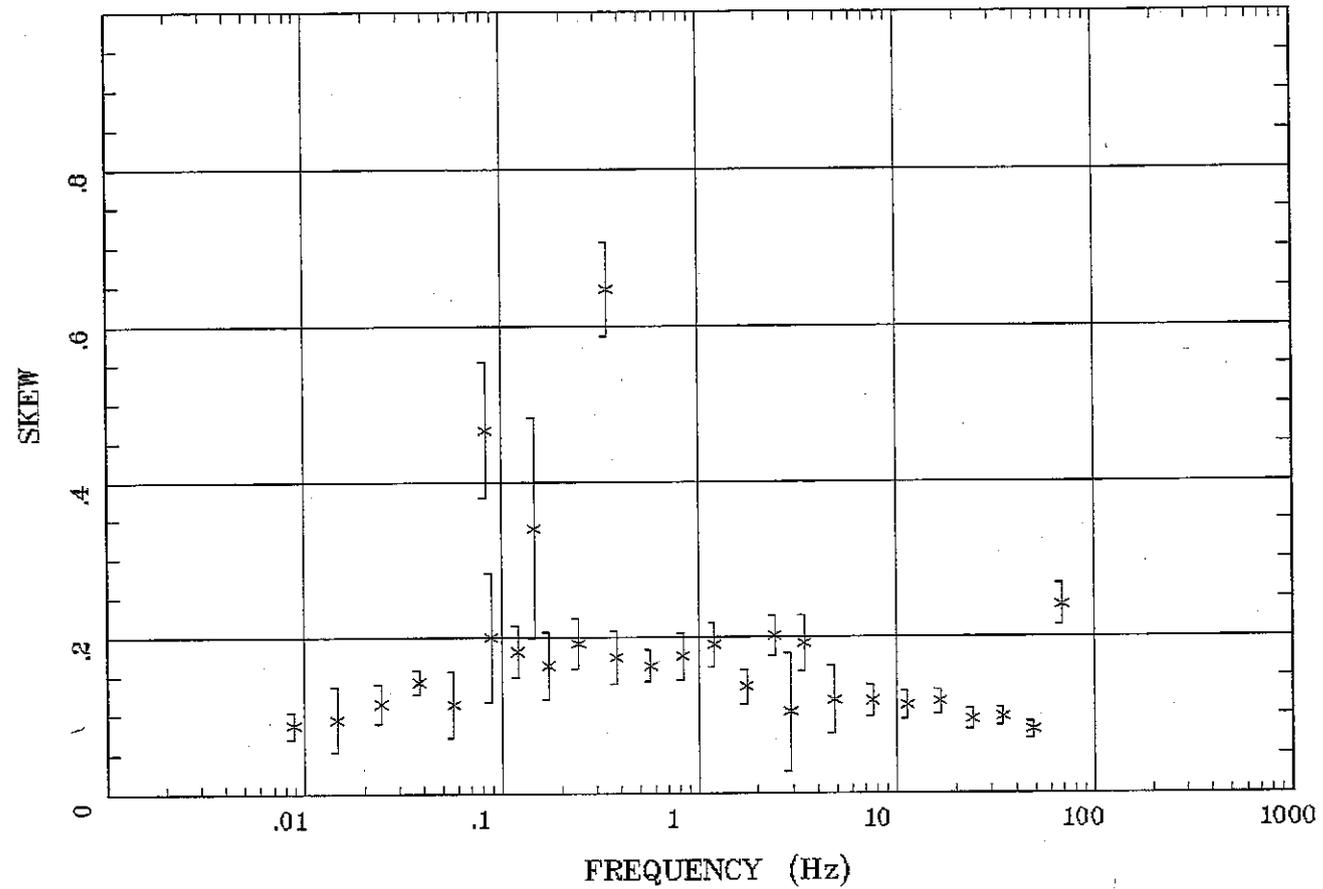
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Station 2

IMPEDANCE SKEW

Wheeler Peak Quad, NM



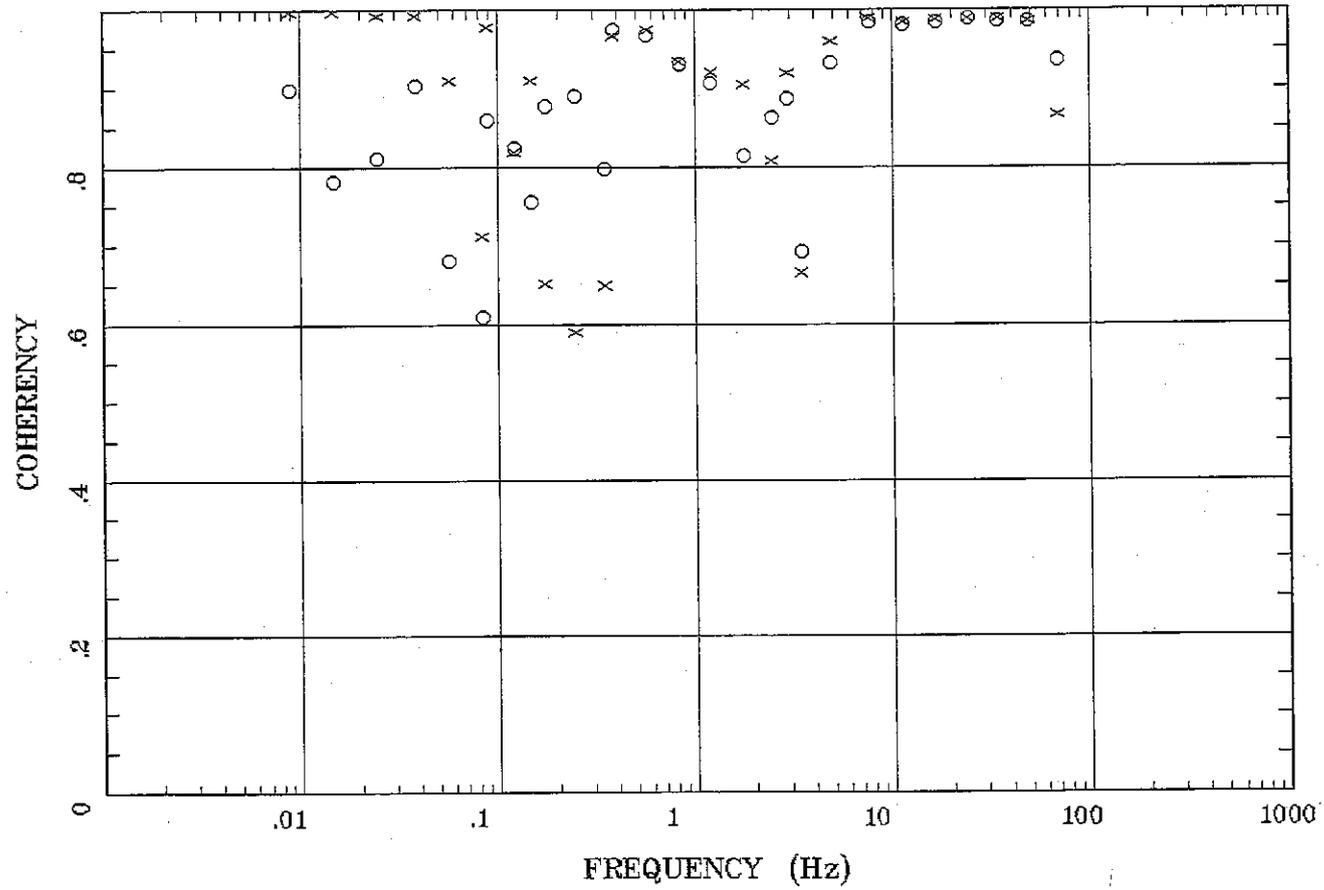
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Survey Co:USGS

Rotation:
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Station 2

E MULT Coh.

Wheeler Peak Quad, NM



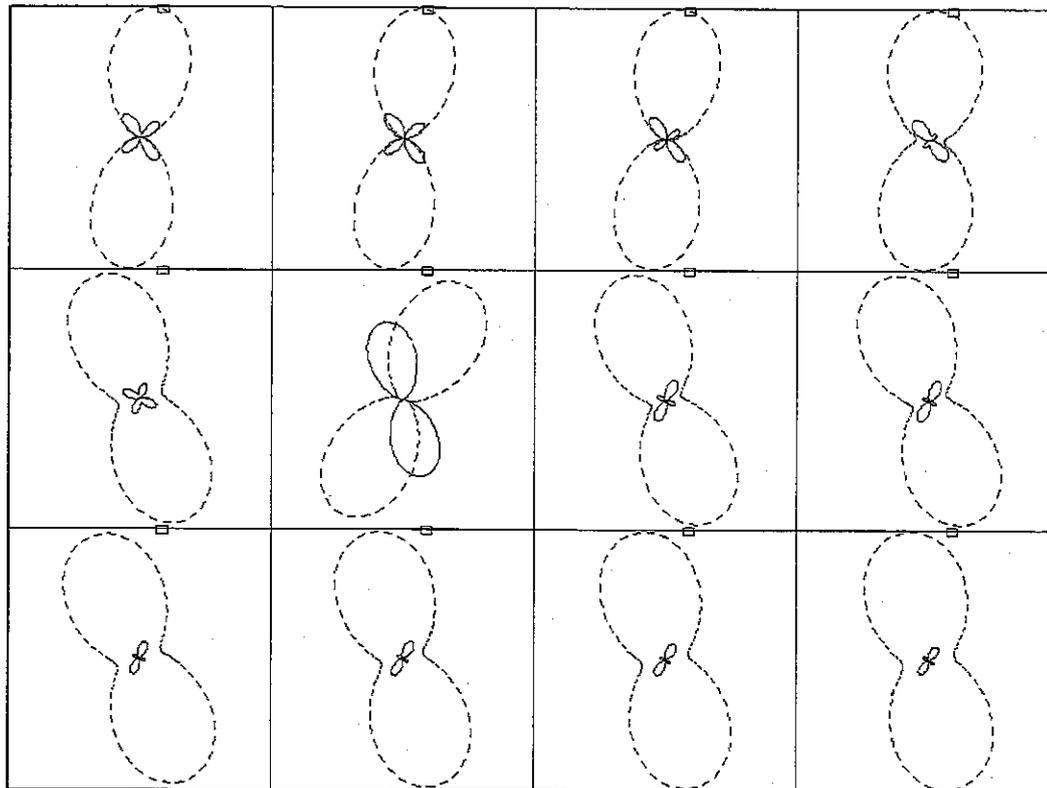
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Survey Co:USGS

Rotation:
Filename: wp02m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 2

POLAR PLOTS

Wheeler Peak Quad, NM



.0088 Hz	.0244 Hz	.0566 Hz	.120 Hz
.172 Hz	.345 Hz	.566 Hz	1.758 Hz
2.930 Hz	7.617 Hz	16.602 Hz	34.975 Hz

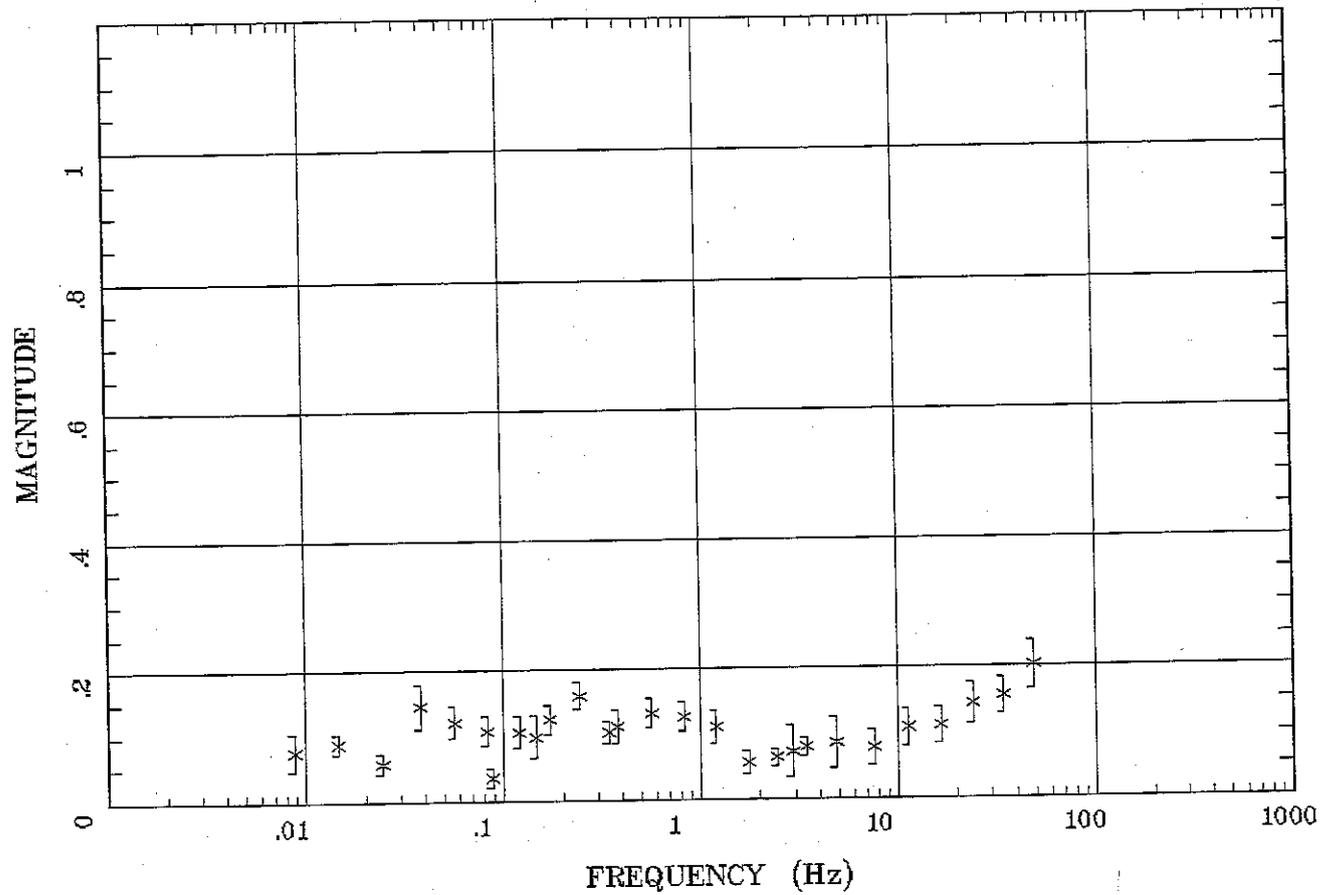
Client:
 Remote: none
 Acquired: 09:3 Jul 24, 2009
 Survey Co:USGS

Rotation:
 Filename: wp02m3.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:51 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 2

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



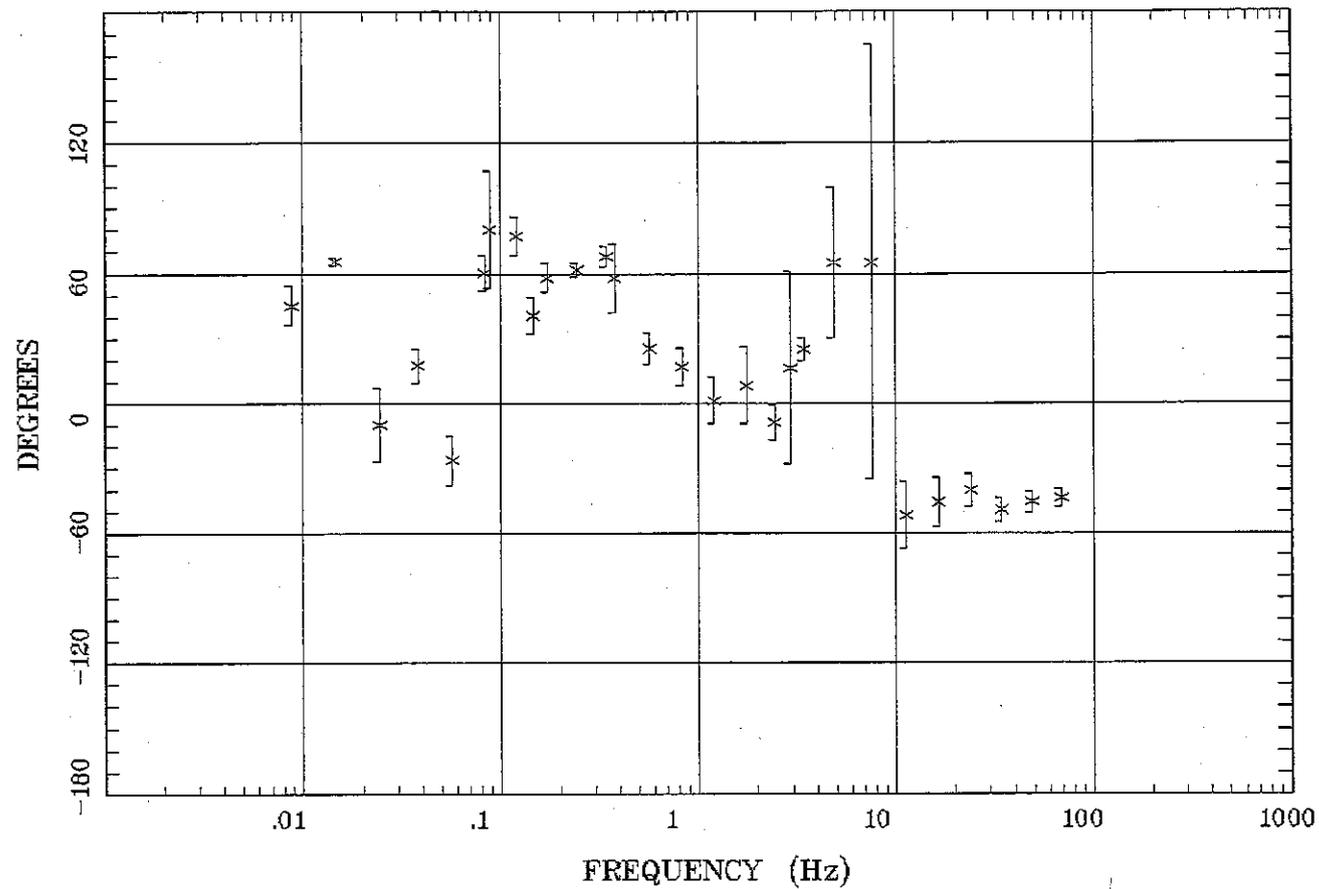
Client:
Remote: none
Acquired: 09:3 Jul 24, 2009
Survey Co:USGS

Rotation:
Filename: wp02m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 2

TIPPER STRIKE

Wheeler Peak Quad, NM



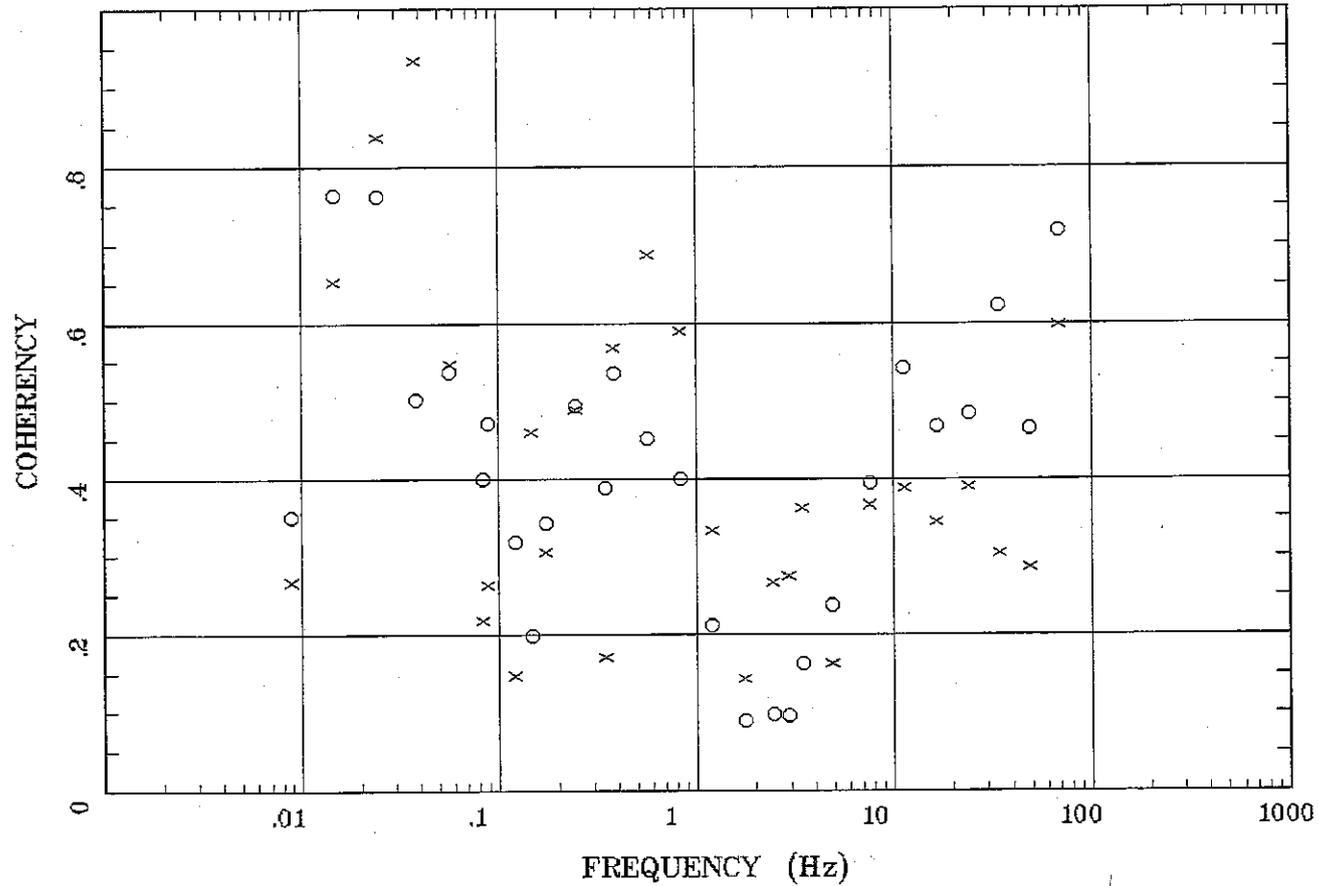
Client:
Remote: none
Acquired: 09:3 Jul 24, 2009
Survey Co:USGS

Rotation:
Filename: wp02m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 2

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



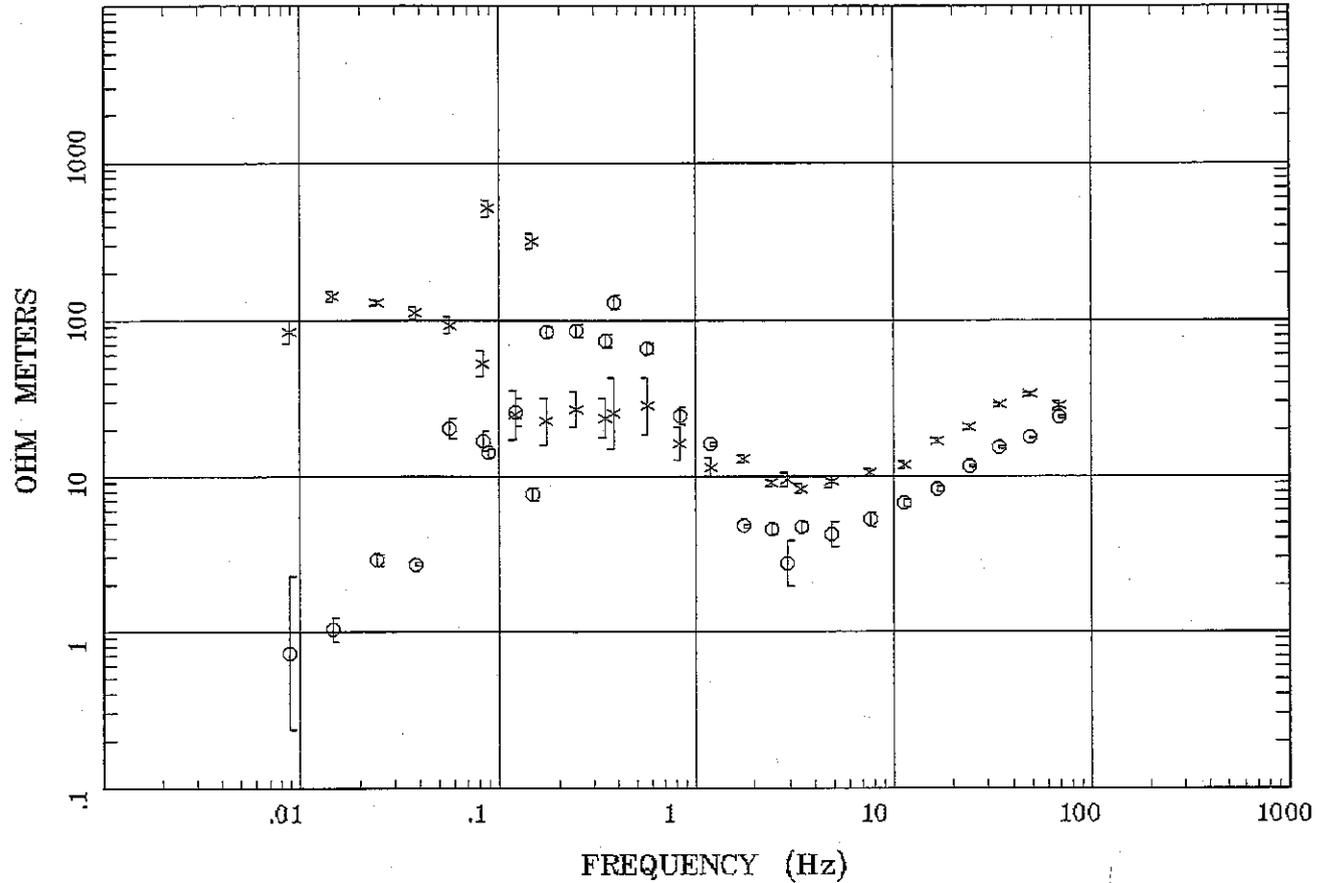
Client:
 Remote: none
 Acquired: 09:3 Jul 24, 2009
 Survey Co:USGS

Rotation:
 Filename: wp02m3.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:51 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 3

APPARENT RESISTIVITY

Wheeler Peak Quad, NM



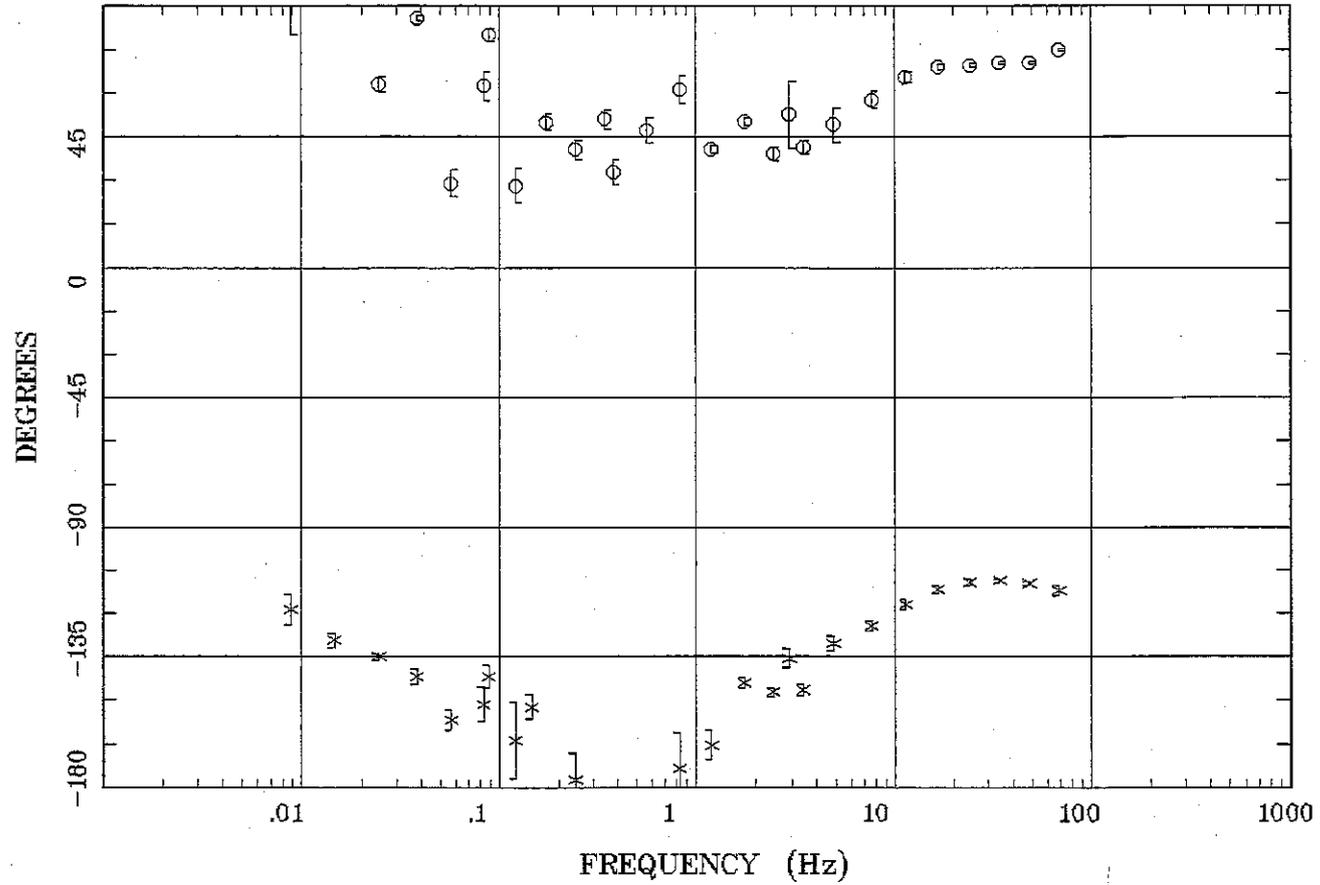
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 3

IMPEDANCE PHASE

Wheeler Peak Quad, NM



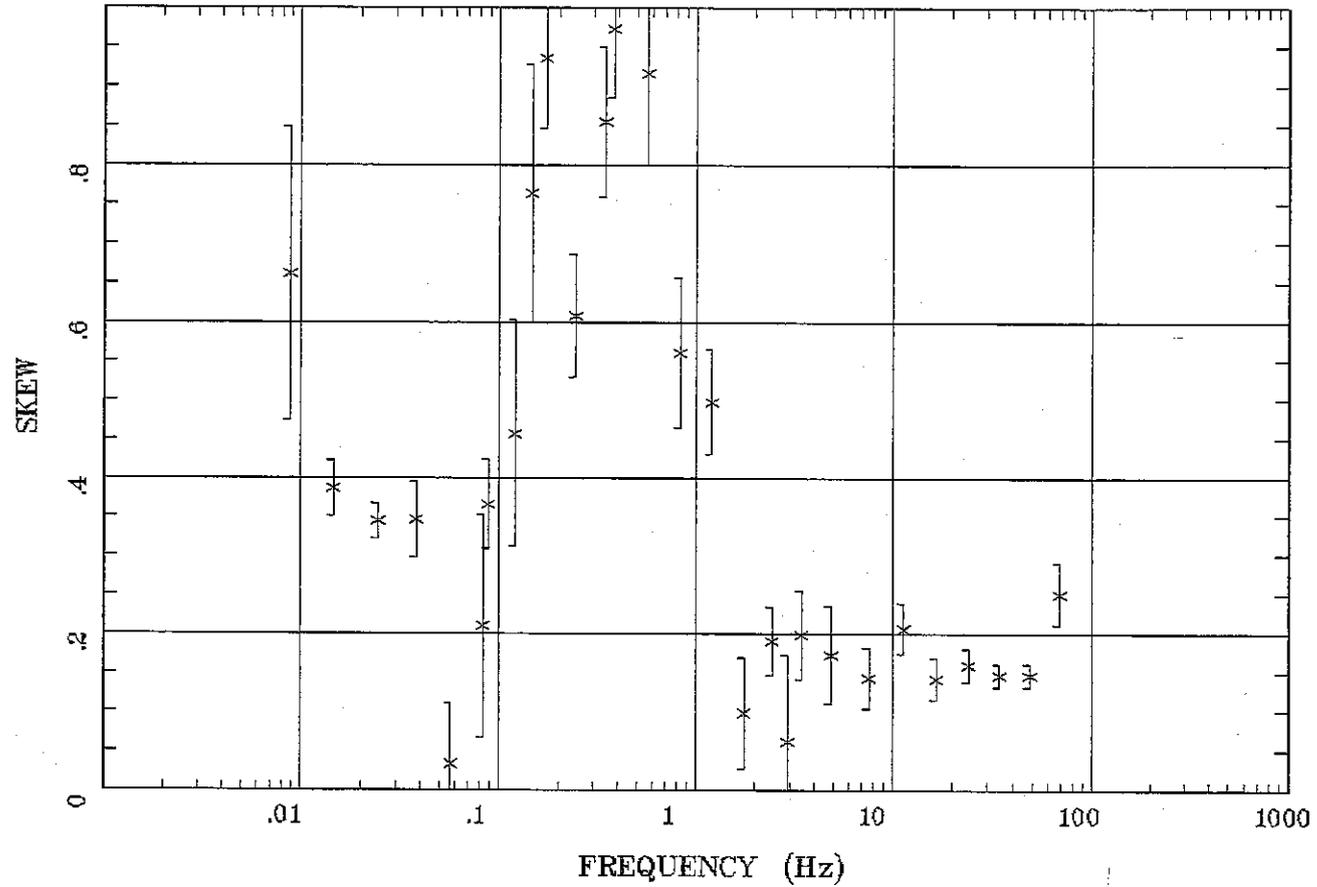
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 3

IMPEDANCE SKEW

Wheeler Peak Quad, NM



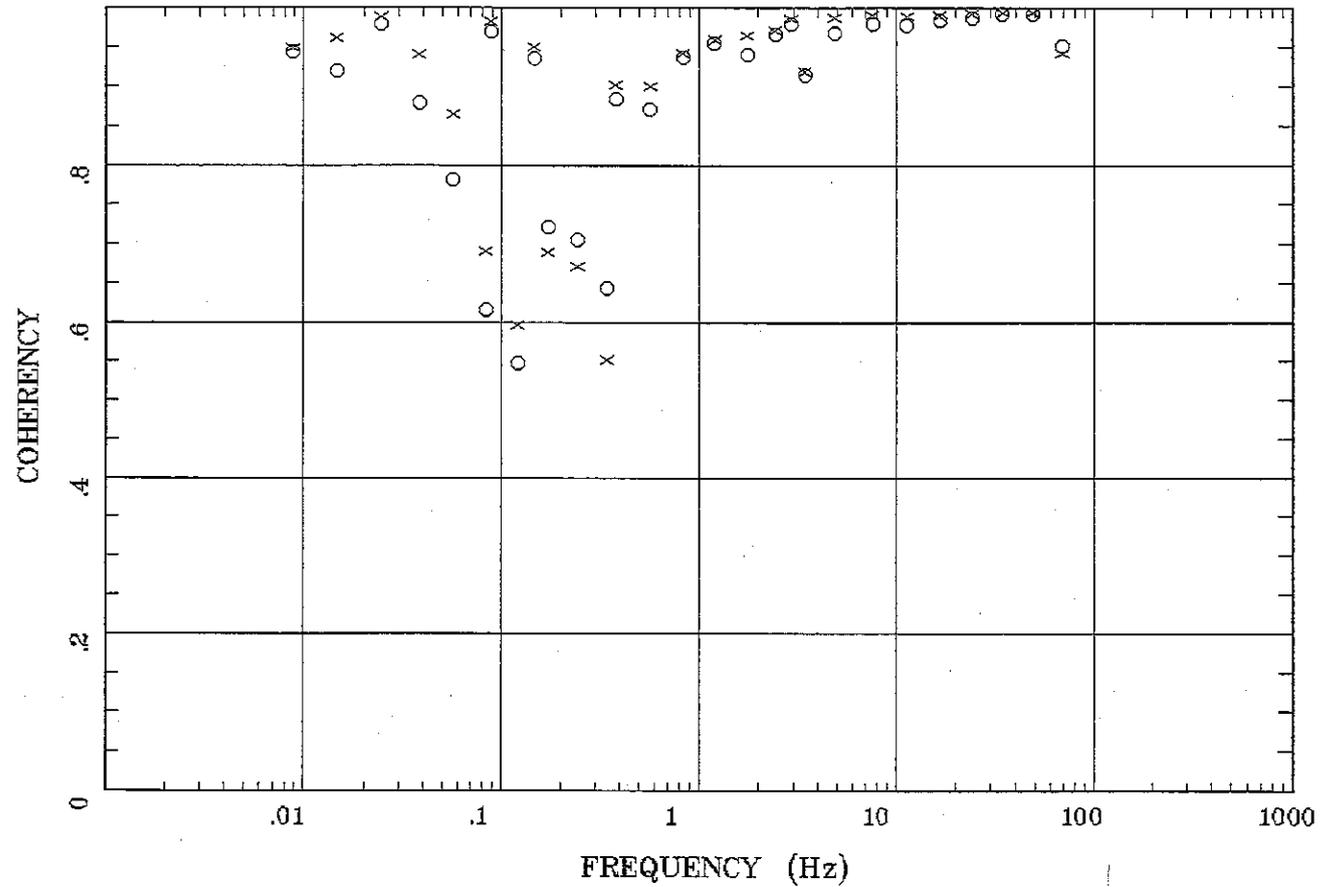
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 3

E MULT Coh.

Wheeler Peak Quad, NM



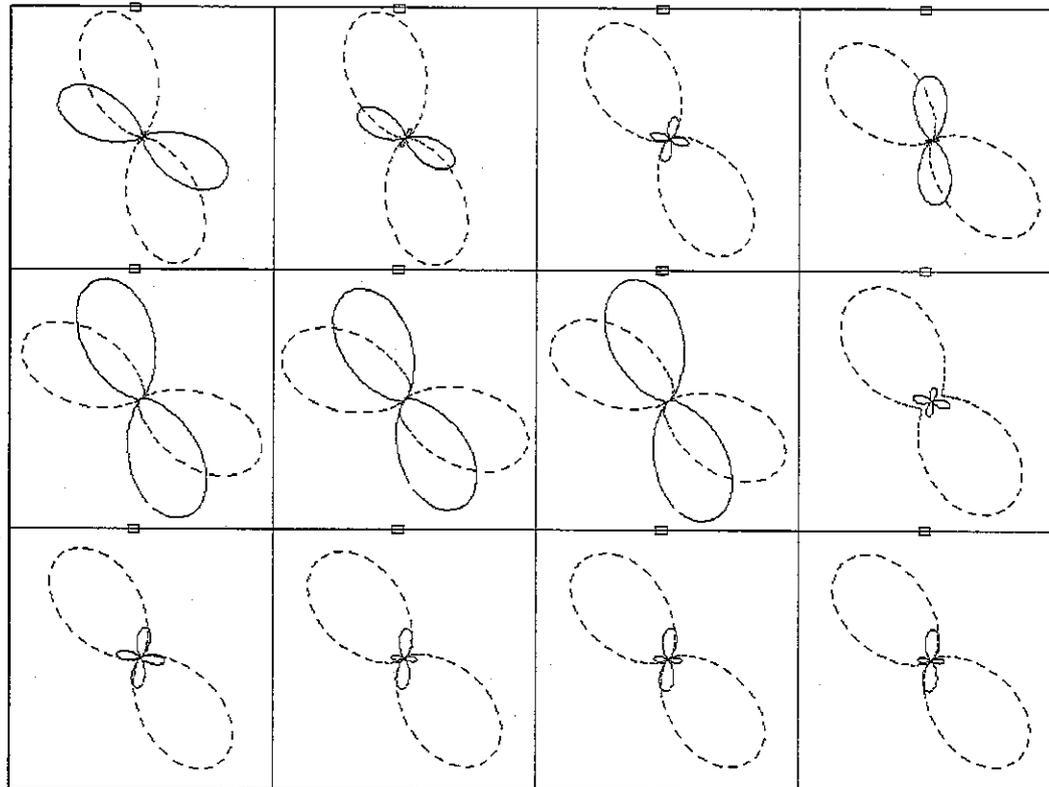
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 3

POLAR PLOTS

Wheeler Peak Quad, NM



.0088 Hz	.0244 Hz	.0566 Hz	.120 Hz
.172 Hz	.345 Hz	.566 Hz	1.758 Hz
2.930 Hz	7.617 Hz	16.602 Hz	34.375 Hz

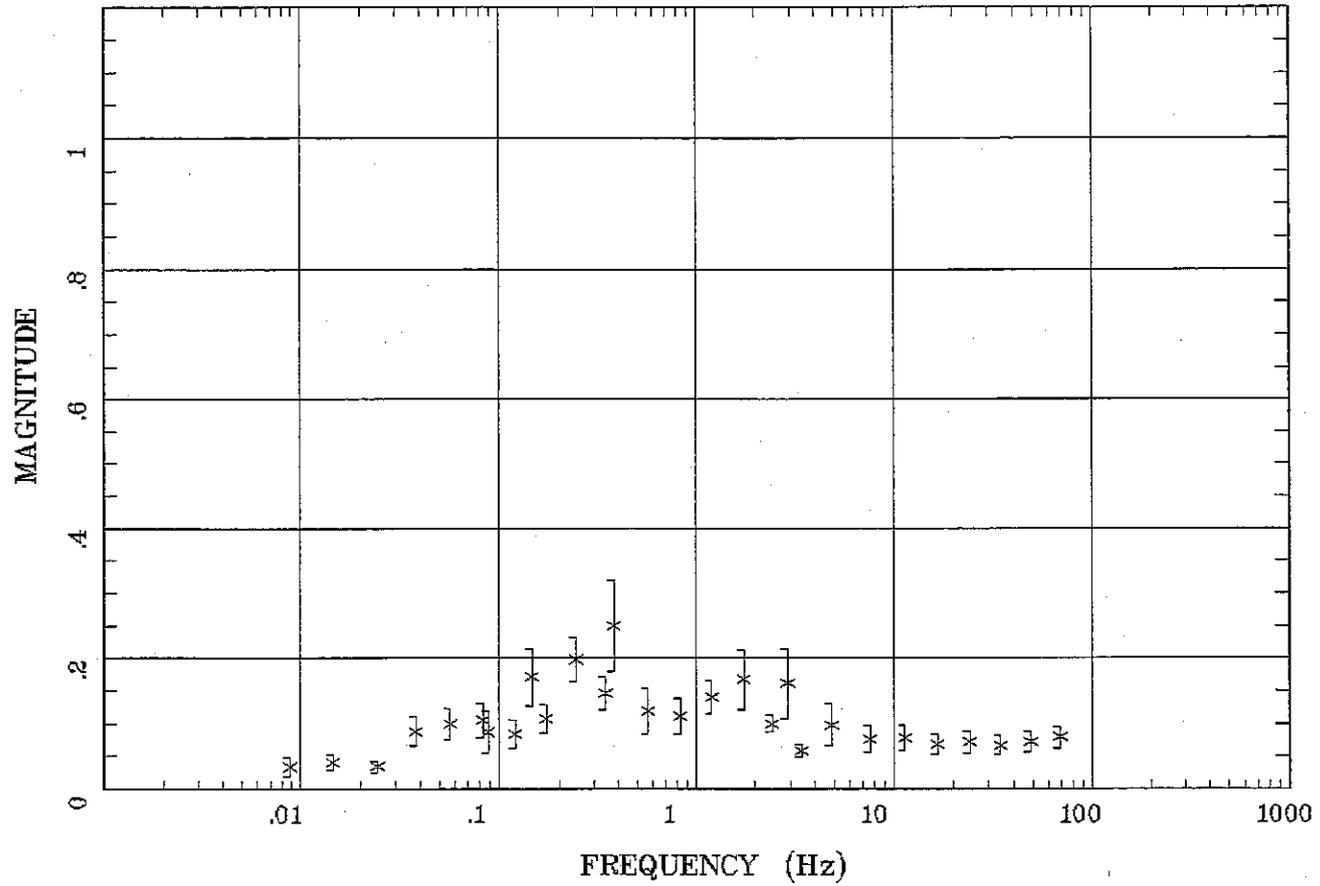
Client:
 Remote: none
 Acquired: 09:5 Jul 22, 2009
 Survey Co:USGS

Rotation:
 Filename: wp03m4.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:51 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 3

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



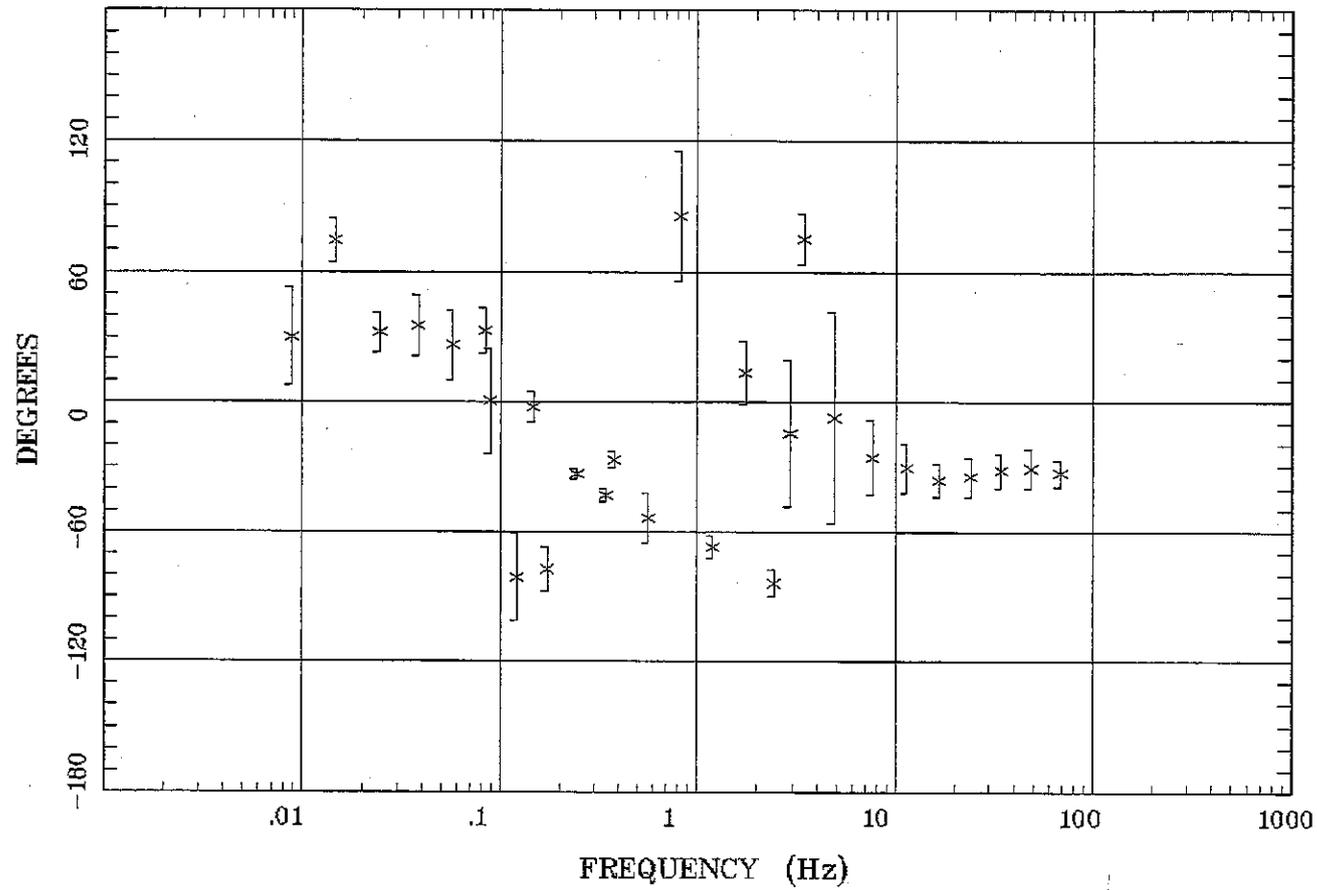
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 3

TIPPER STRIKE

Wheeler Peak Quad, NM



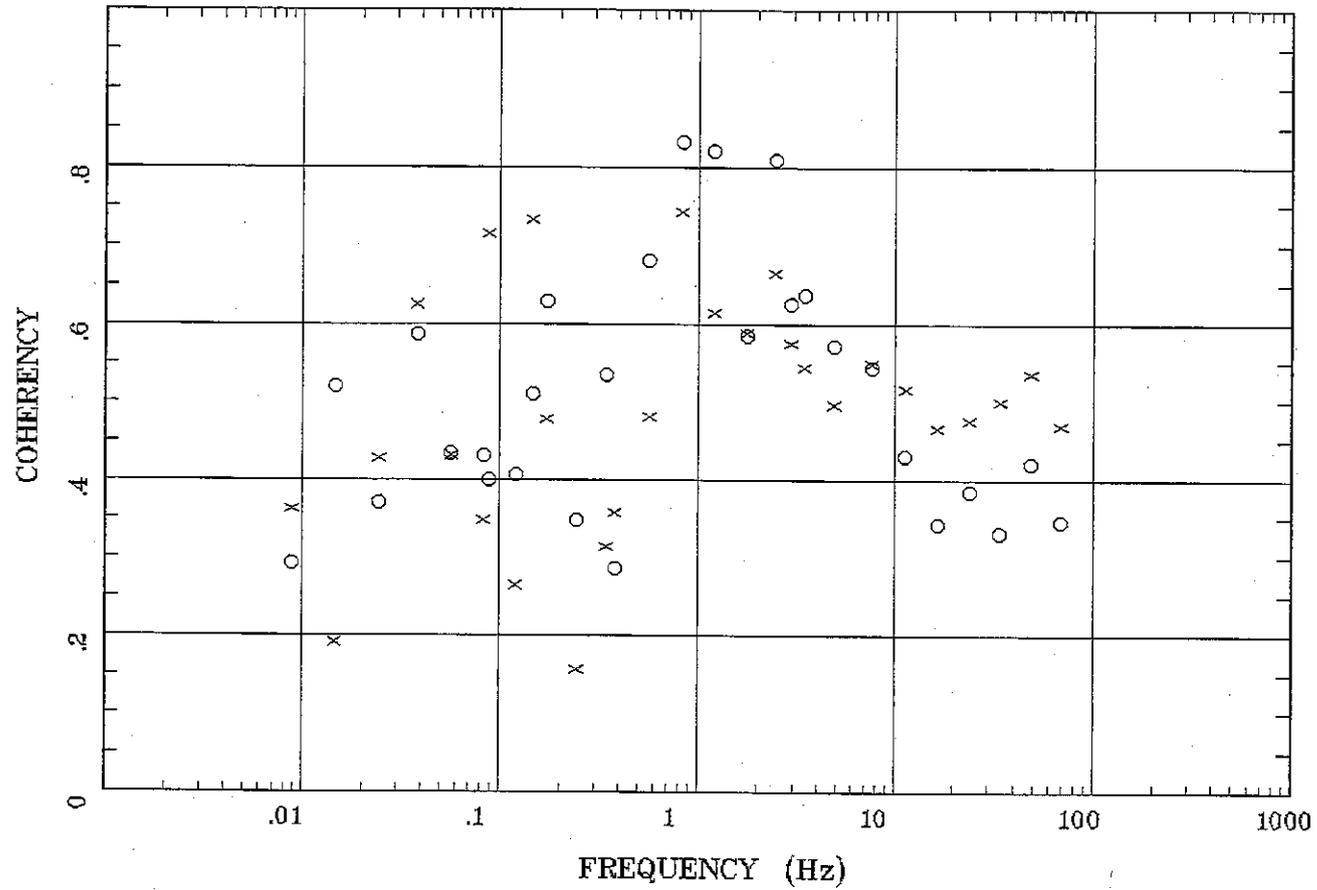
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 3

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



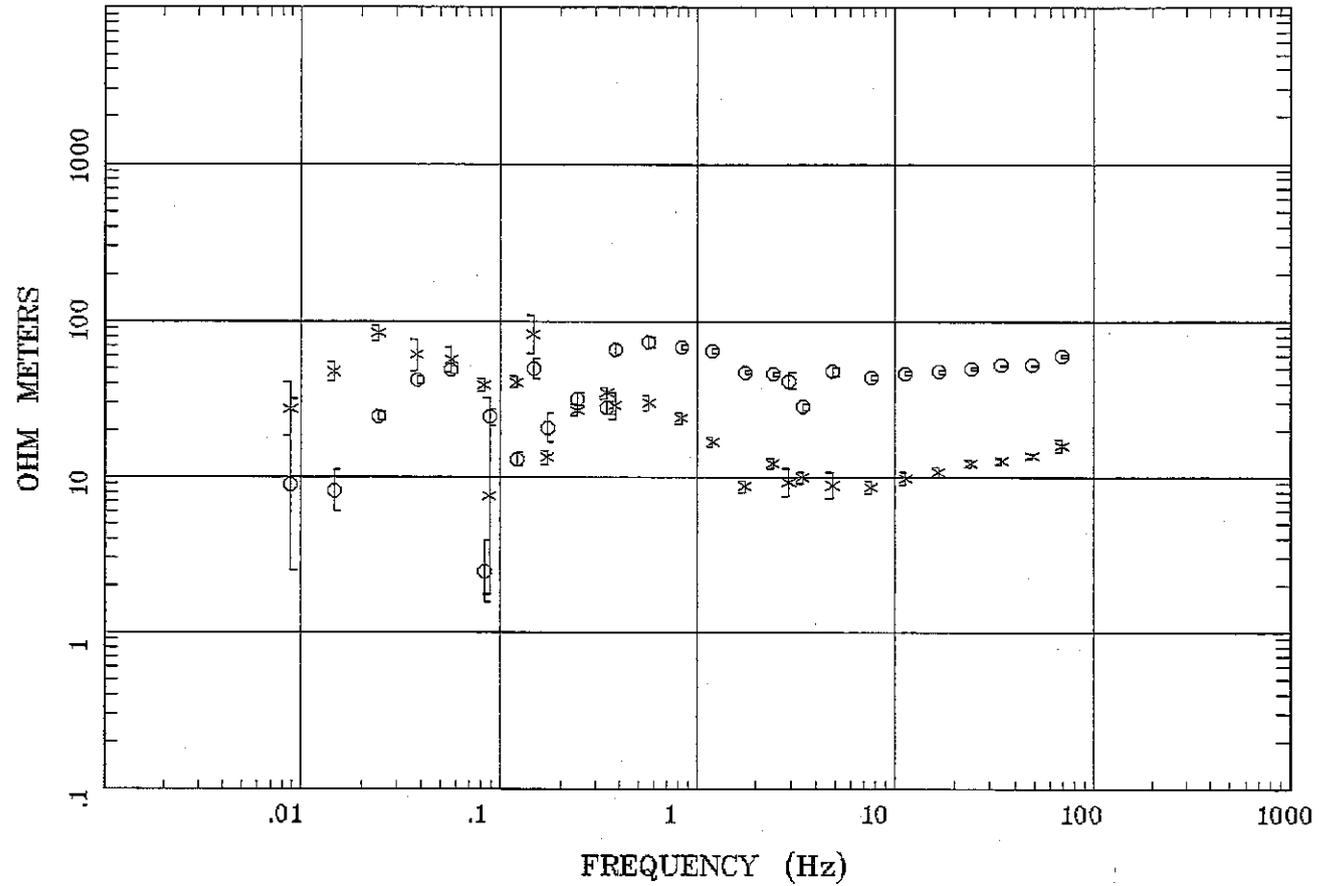
Client:
Remote: none
Acquired: 09:5 Jul 22, 2009
Survey Co:USGS

Rotation:
Filename: wp03m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:51 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

APPARENT RESISTIVITY

Wheeler Peak Quad, NM



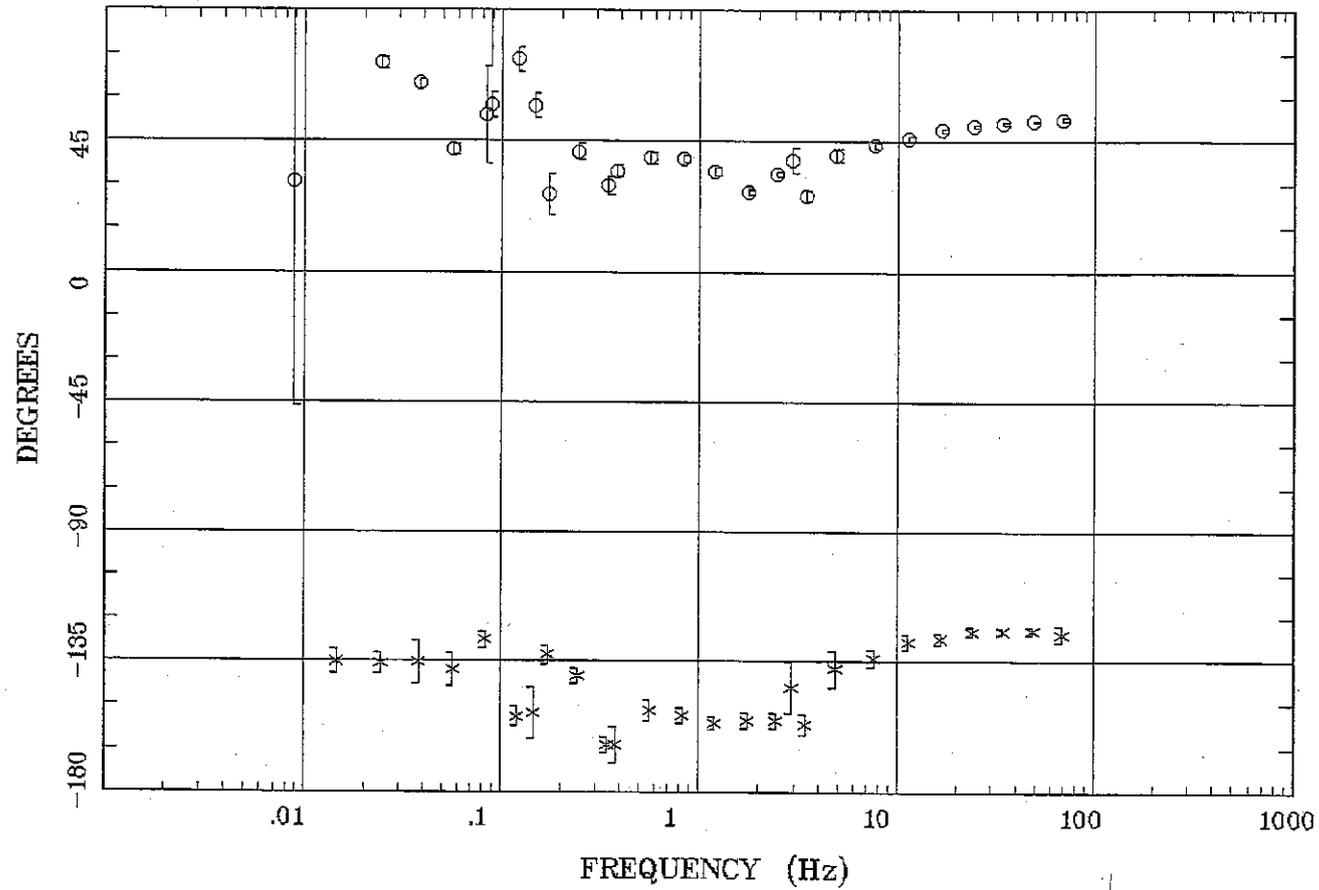
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

IMPEDANCE PHASE

Wheeler Peak Quad, NM



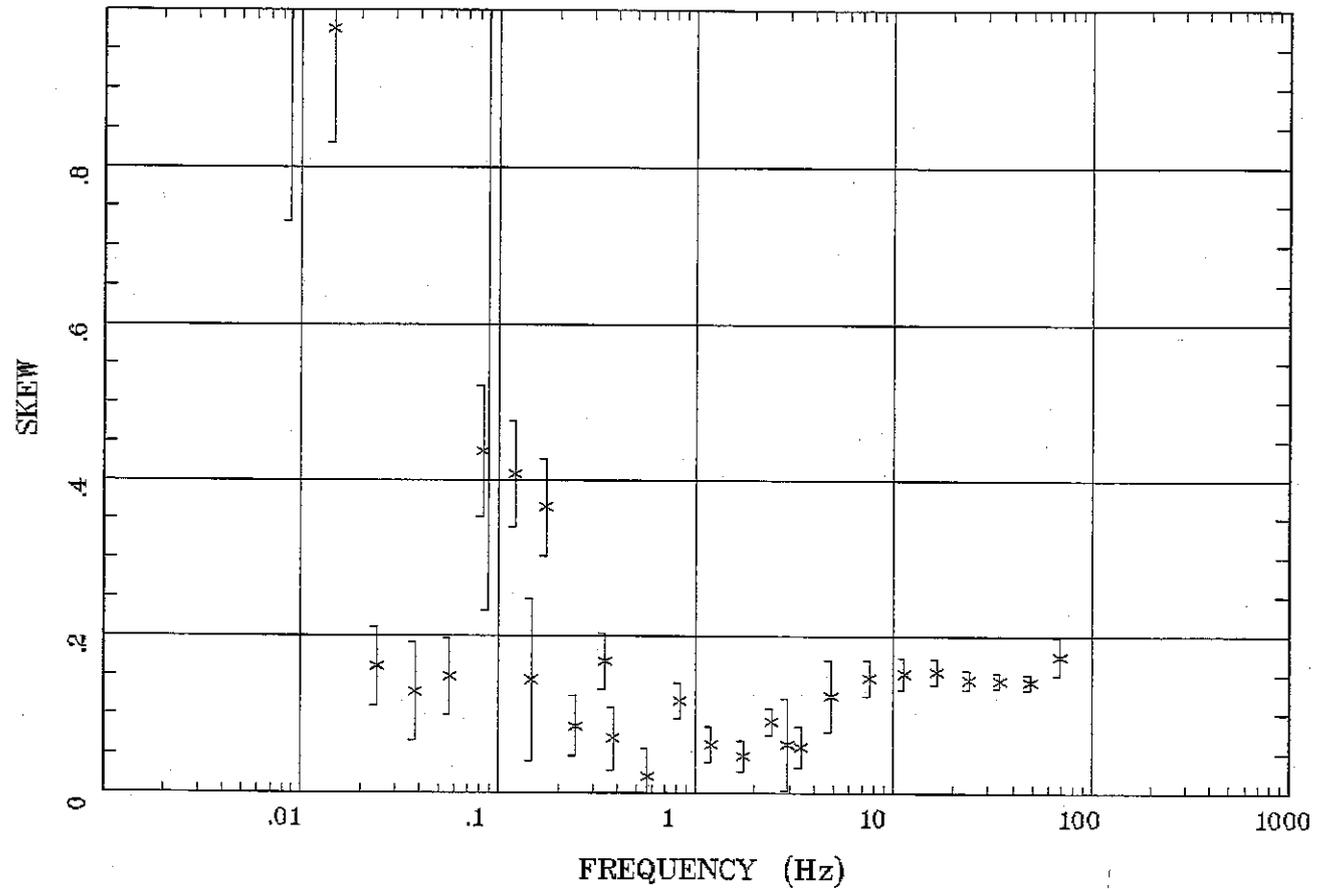
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

IMPEDANCE SKEW

Wheeler Peak Quad, NM



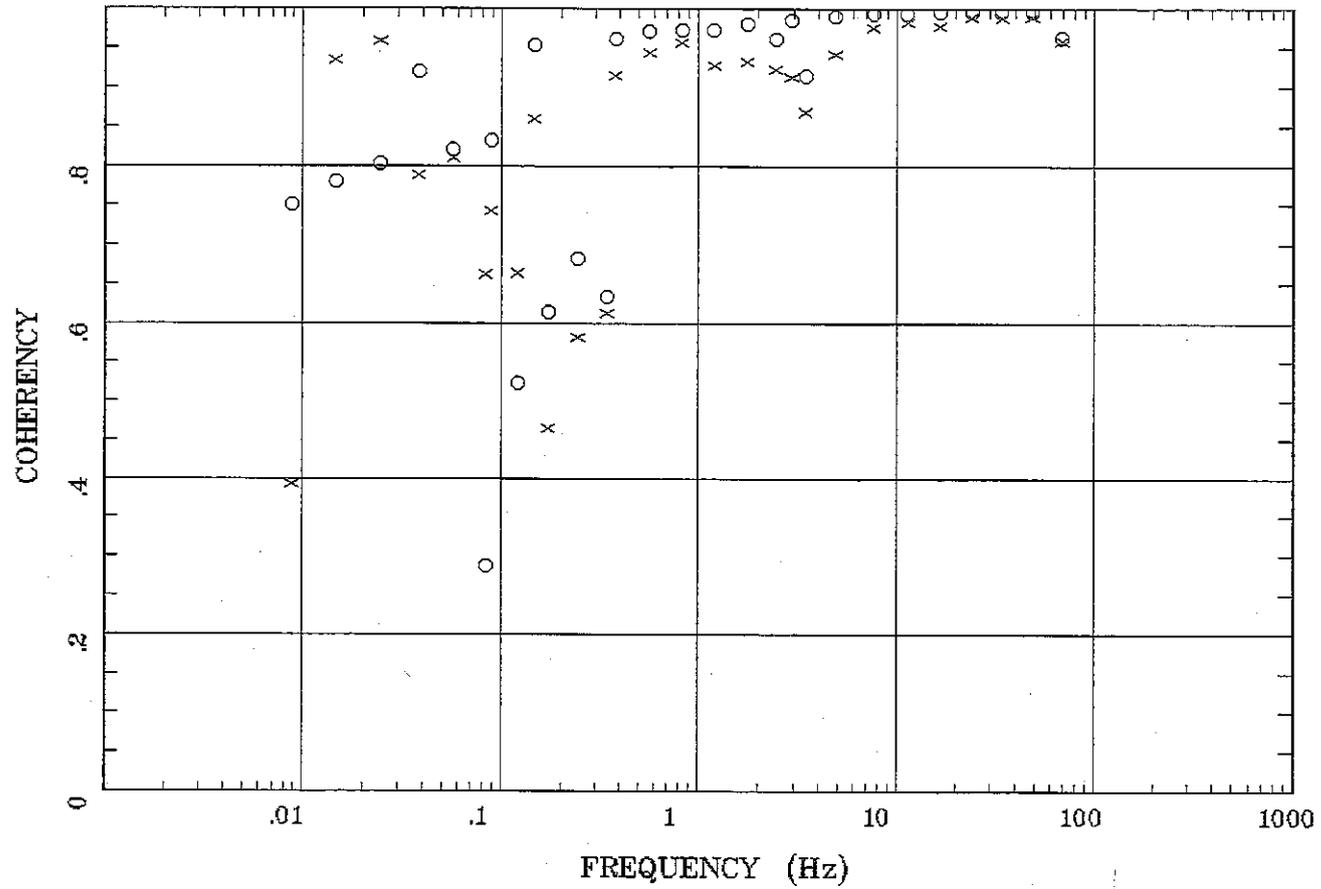
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

E MULT Coh.

Wheeler Peak Quad, NM



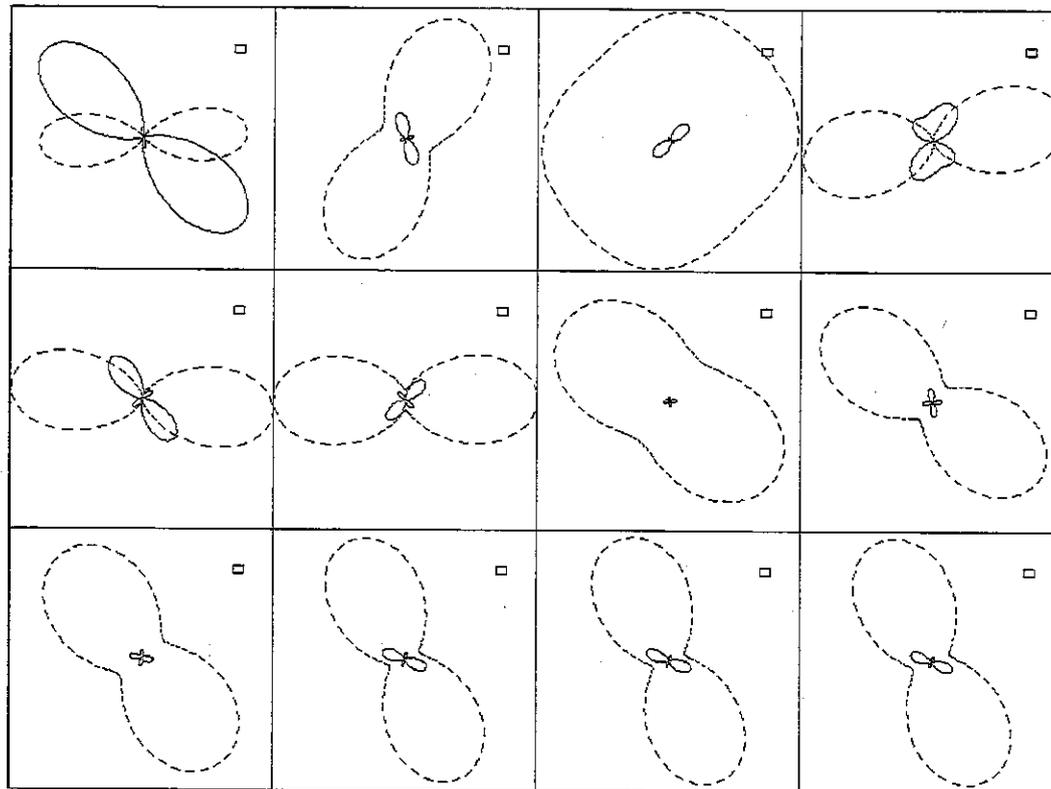
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

POLAR PLOTS

Wheeler Peak Quad, NM



.0088 Hz	.0244 Hz	.0566 Hz	.120 Hz
.172 Hz	.345 Hz	.566 Hz	1.758 Hz
2.930 Hz	7.617 Hz	16.602 Hz	34.875 Hz

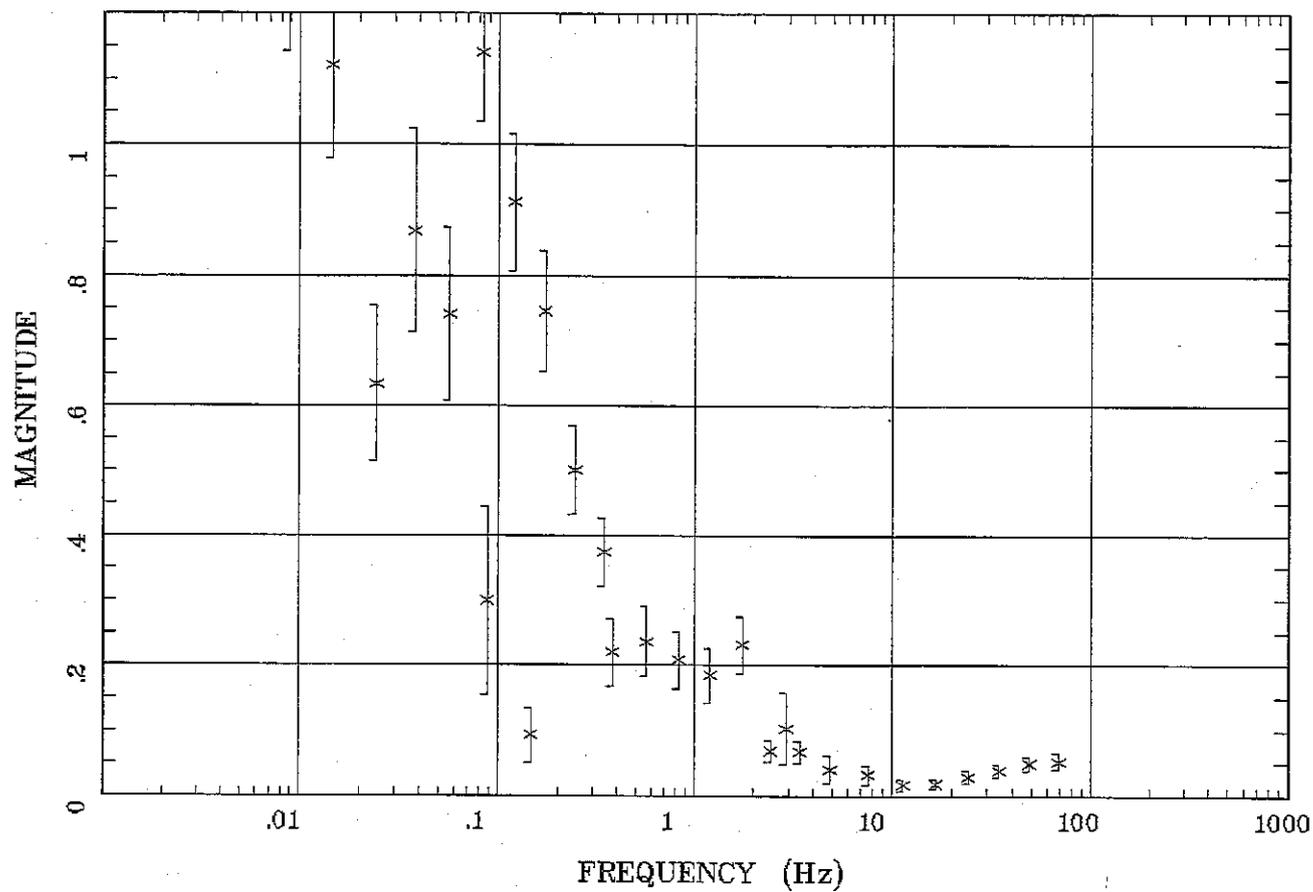
Client:
 Remote: none
 Acquired: 11:0 Jul 25, 2009
 Survey Co:USGS

Rotation:
 Filename: wp04m4.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:53 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 4

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



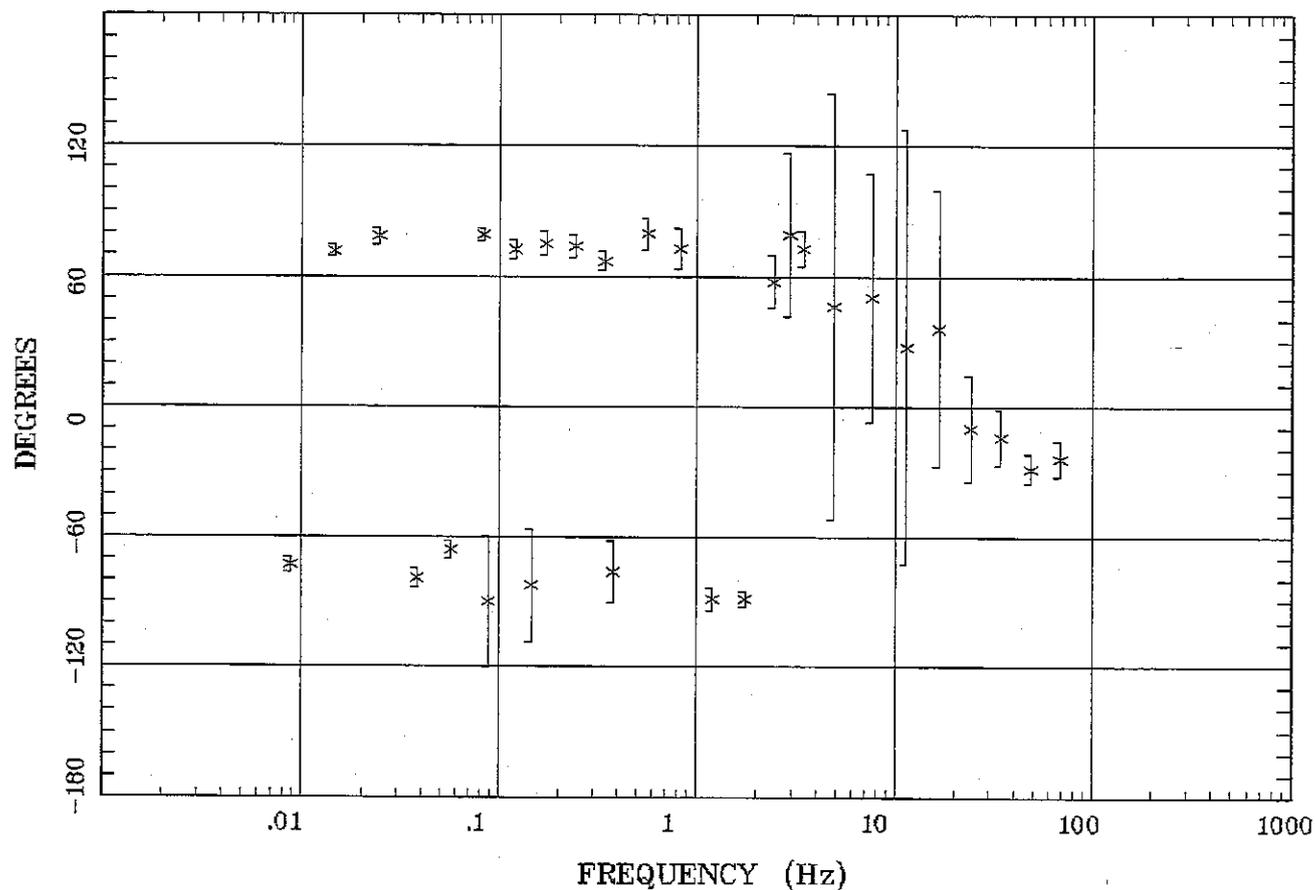
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

TIPPER STRIKE

Wheeler Peak Quad, NM



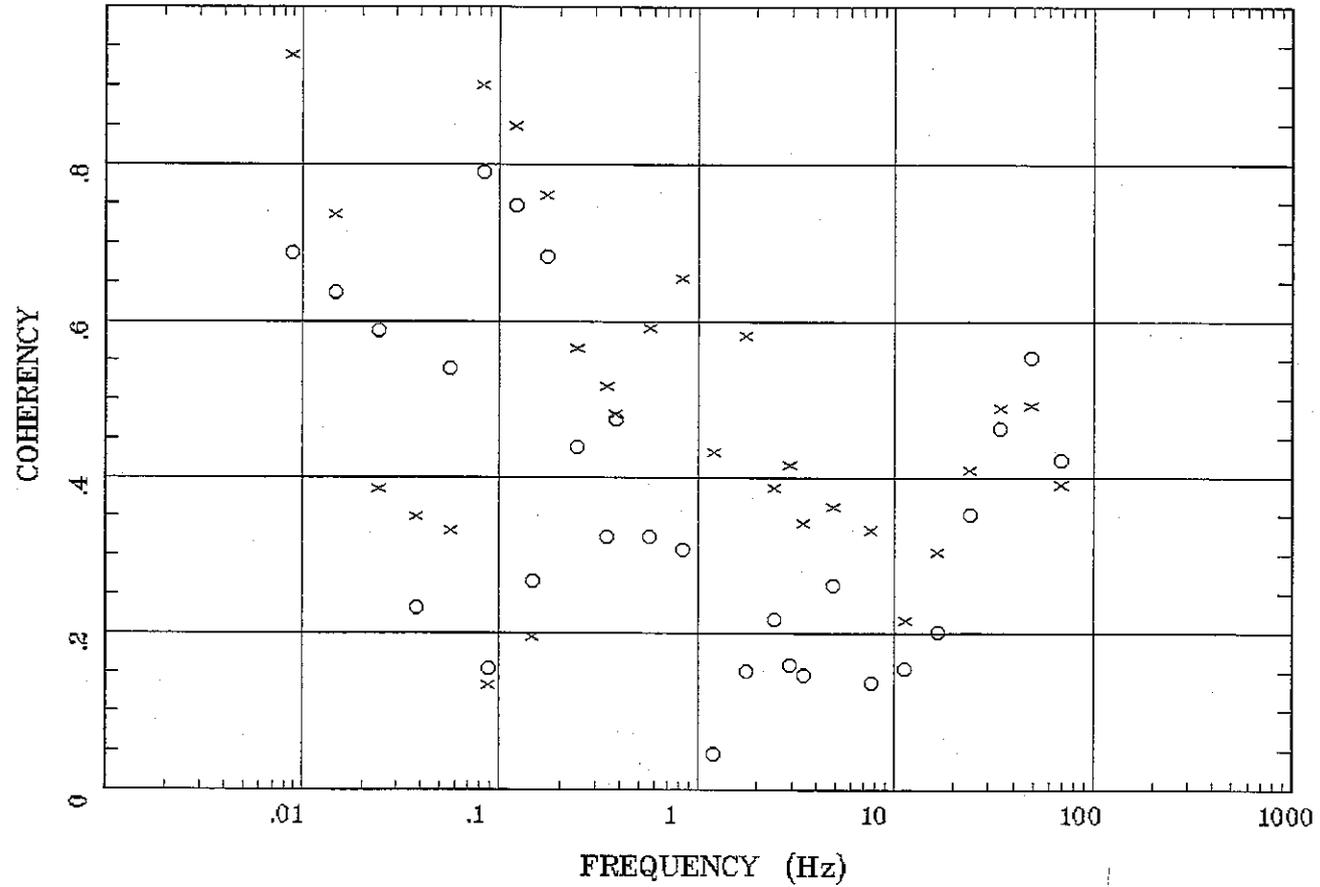
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 4

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



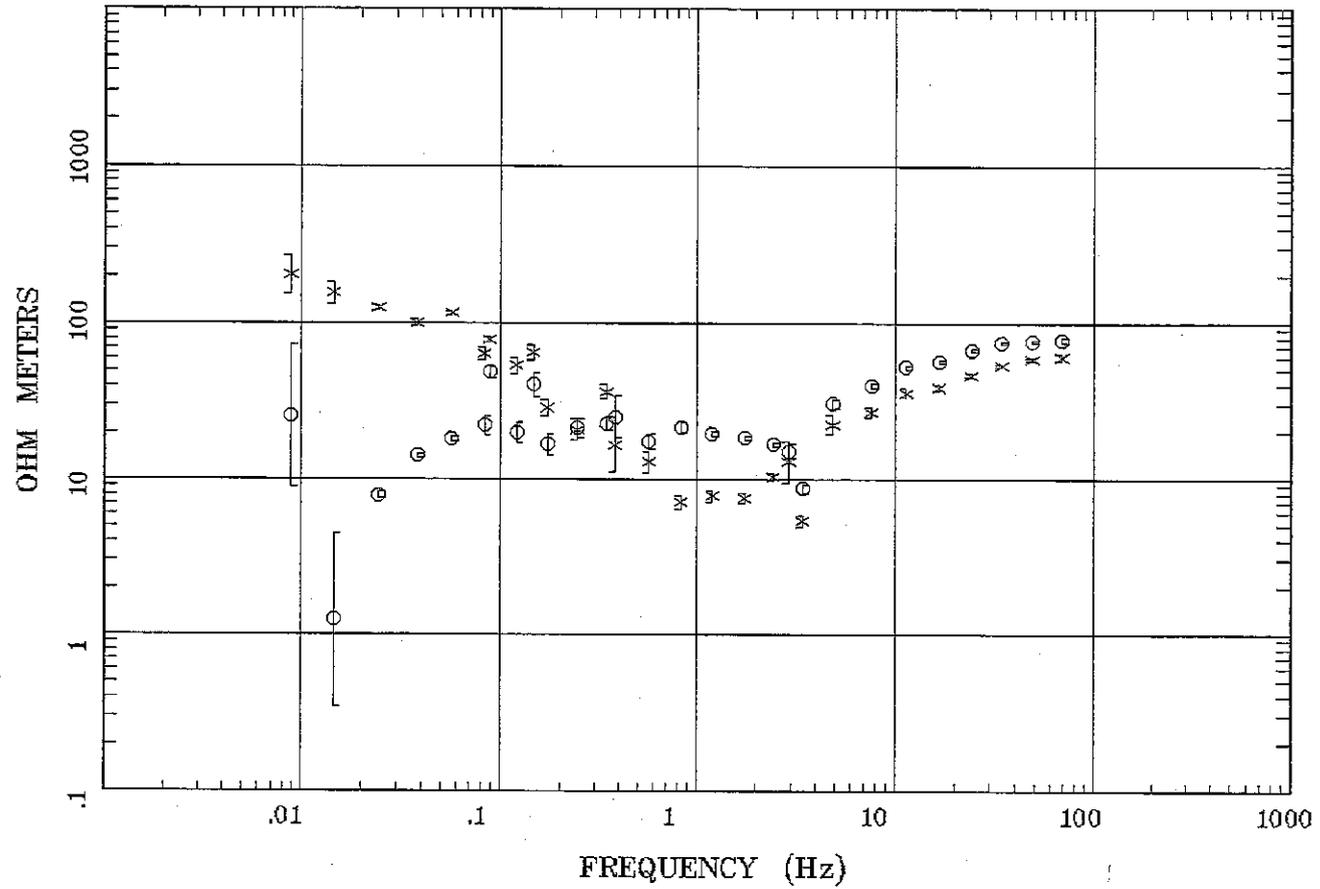
Client:
Remote: none
Acquired: 11:0 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp04m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

APPARENT RESISTIVITY

Wheeler Peak Quad, NM



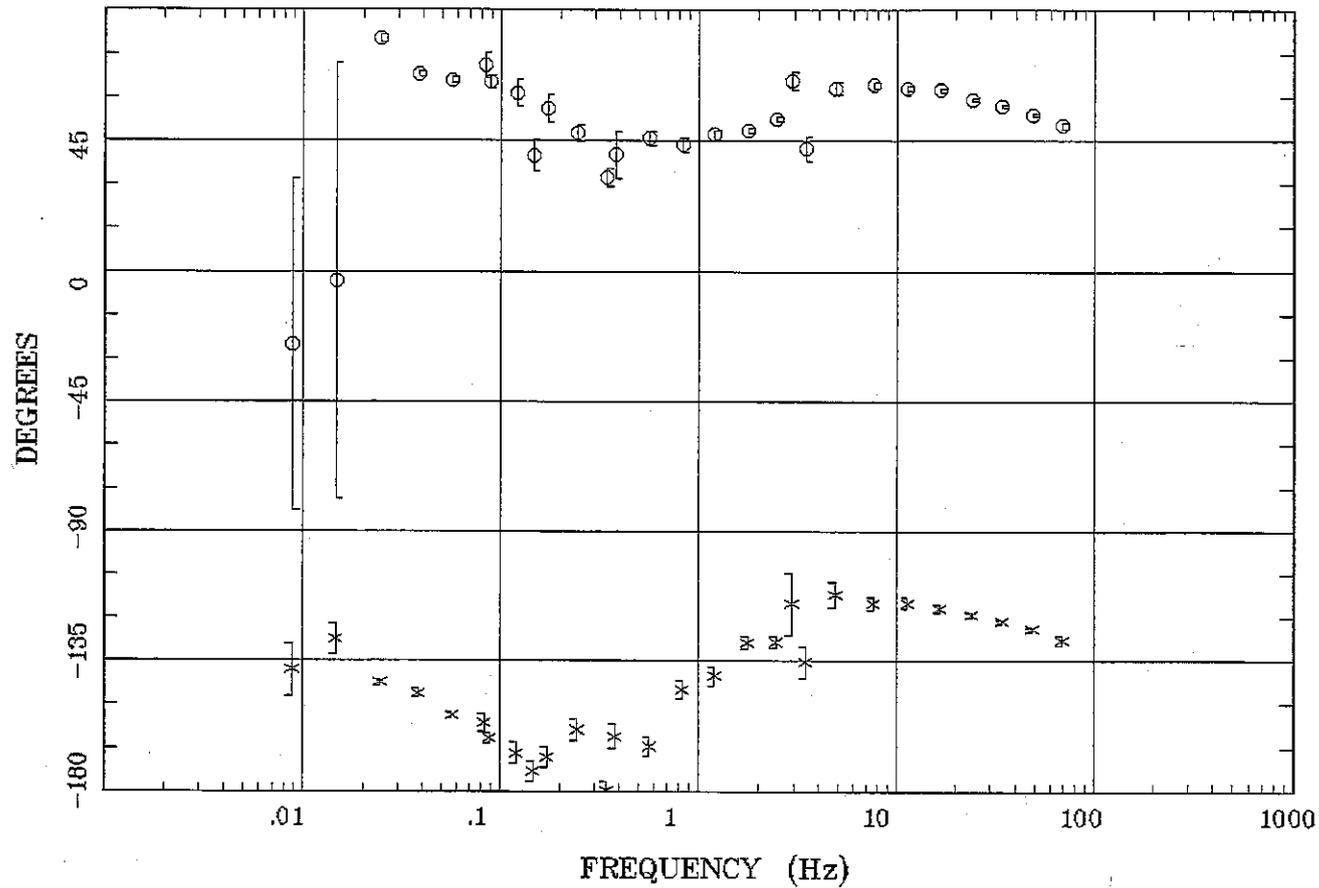
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

IMPEDANCE PHASE

Wheeler Peak Quad, NM



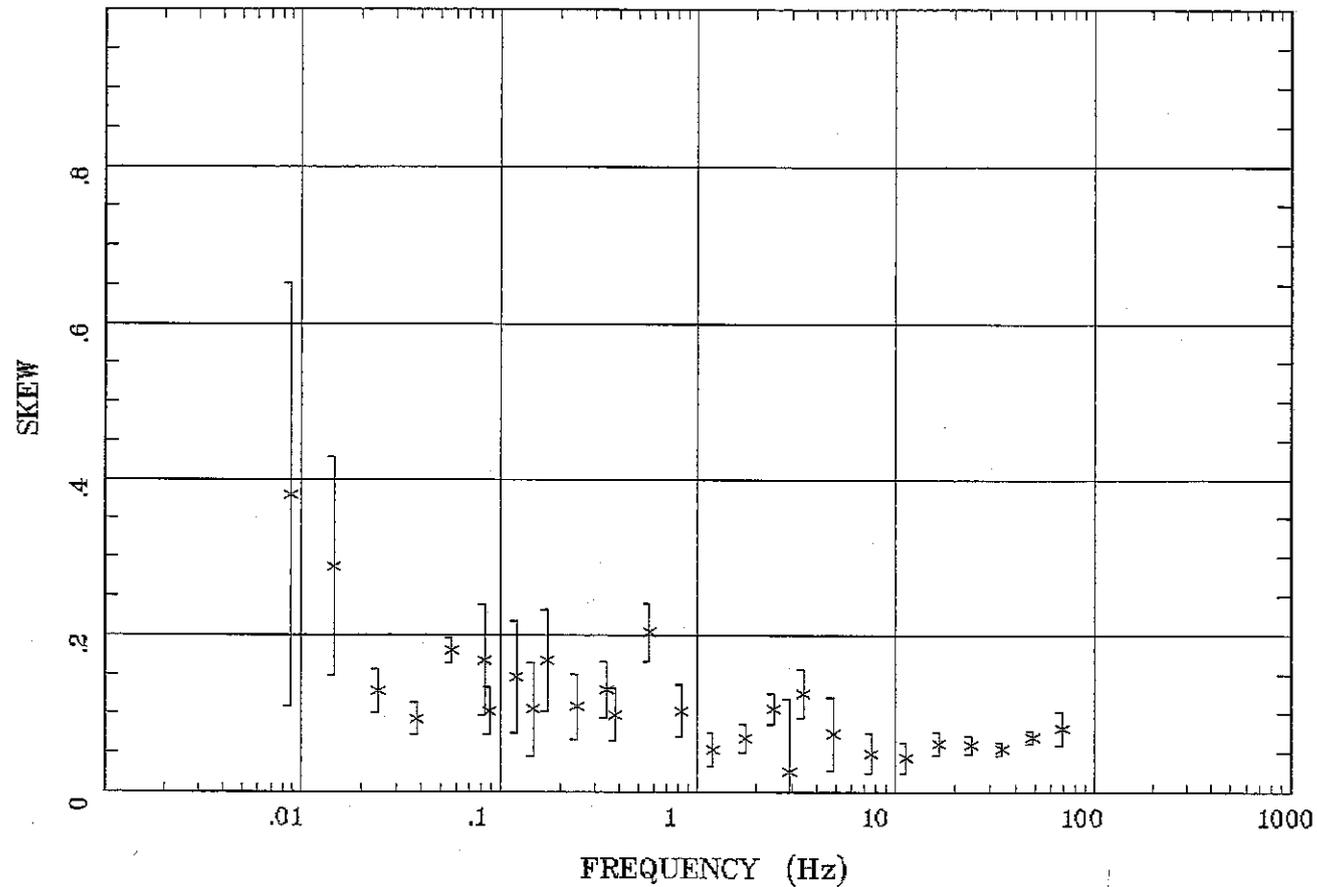
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

IMPEDANCE SKEW

Wheeler Peak Quad, NM



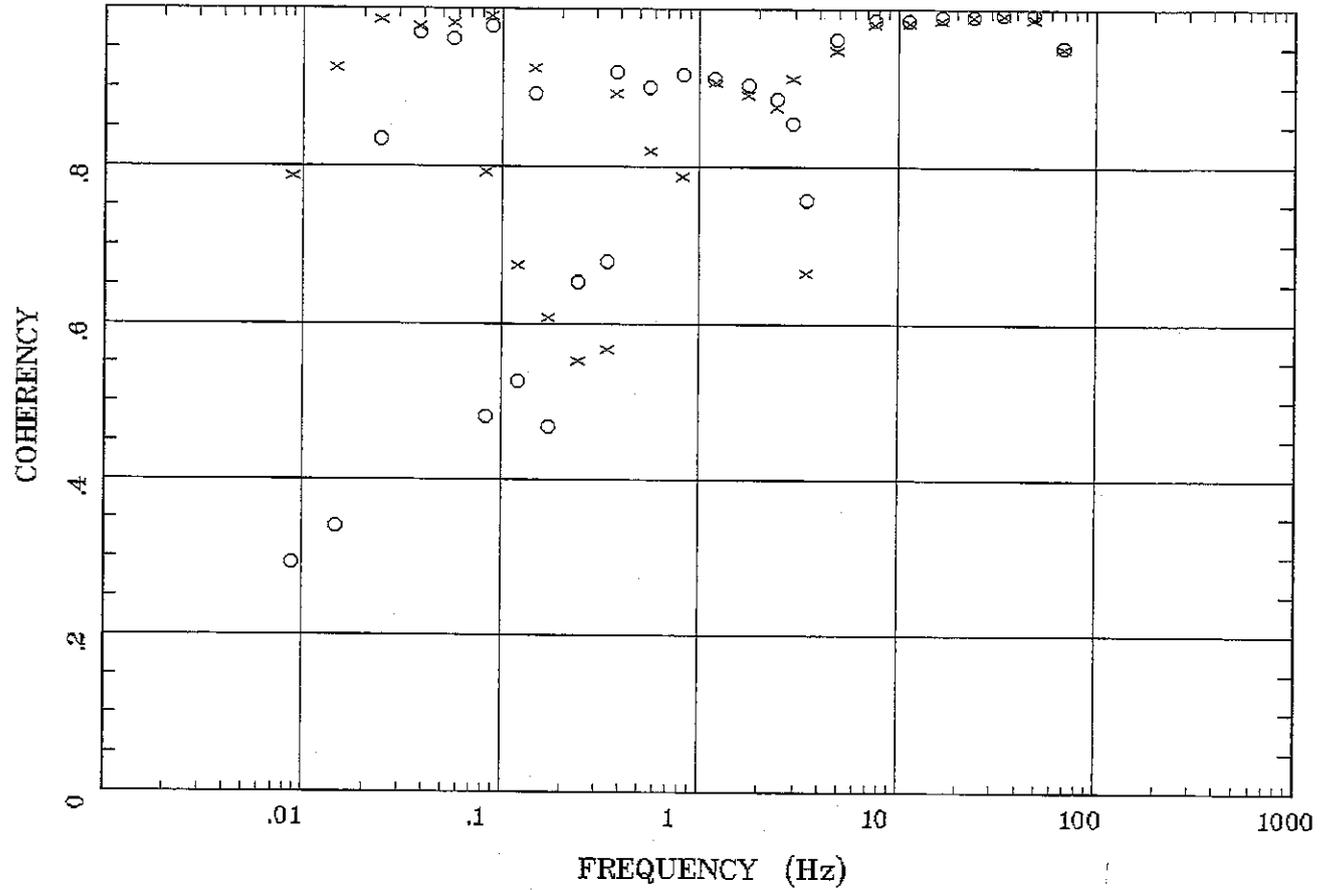
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

E MULT Coh.

Wheeler Peak Quad, NM



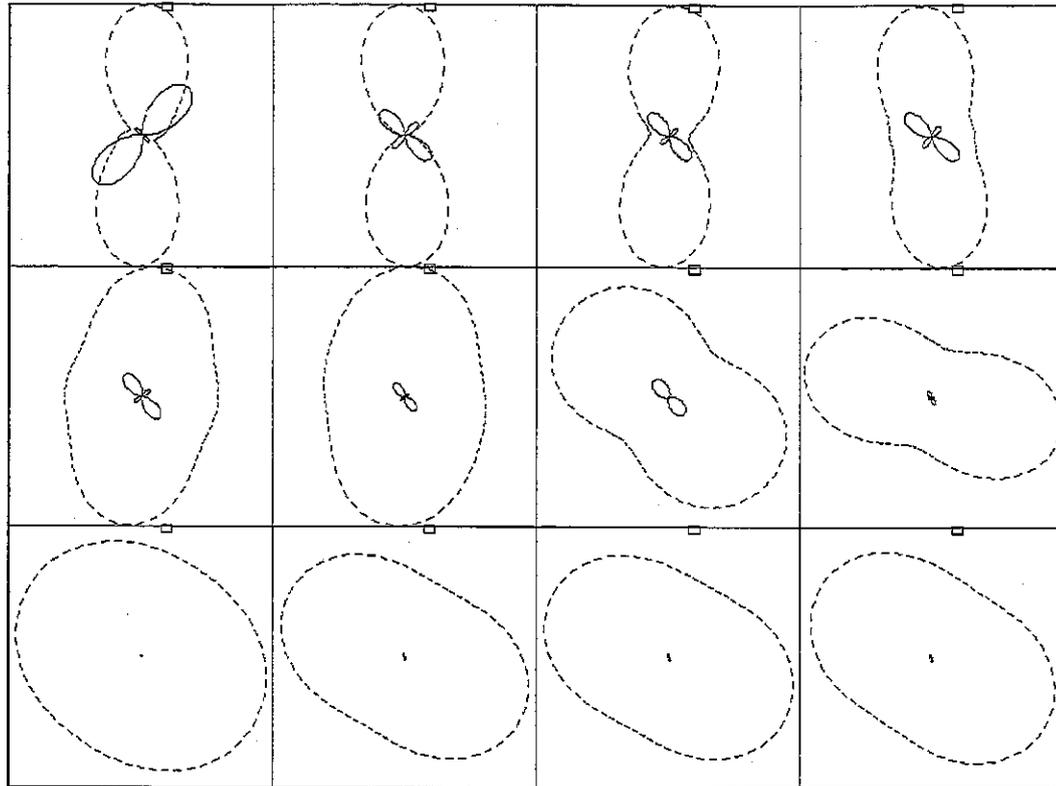
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

POLAR PLOTS

Wheeler Peak Quad, NM



.0088 Hz	.0244 Hz	.0566 Hz	.120 Hz
.172 Hz	.345 Hz	.566 Hz	1.758 Hz
2.930 Hz	7.617 Hz	16.602 Hz	34.375 Hz

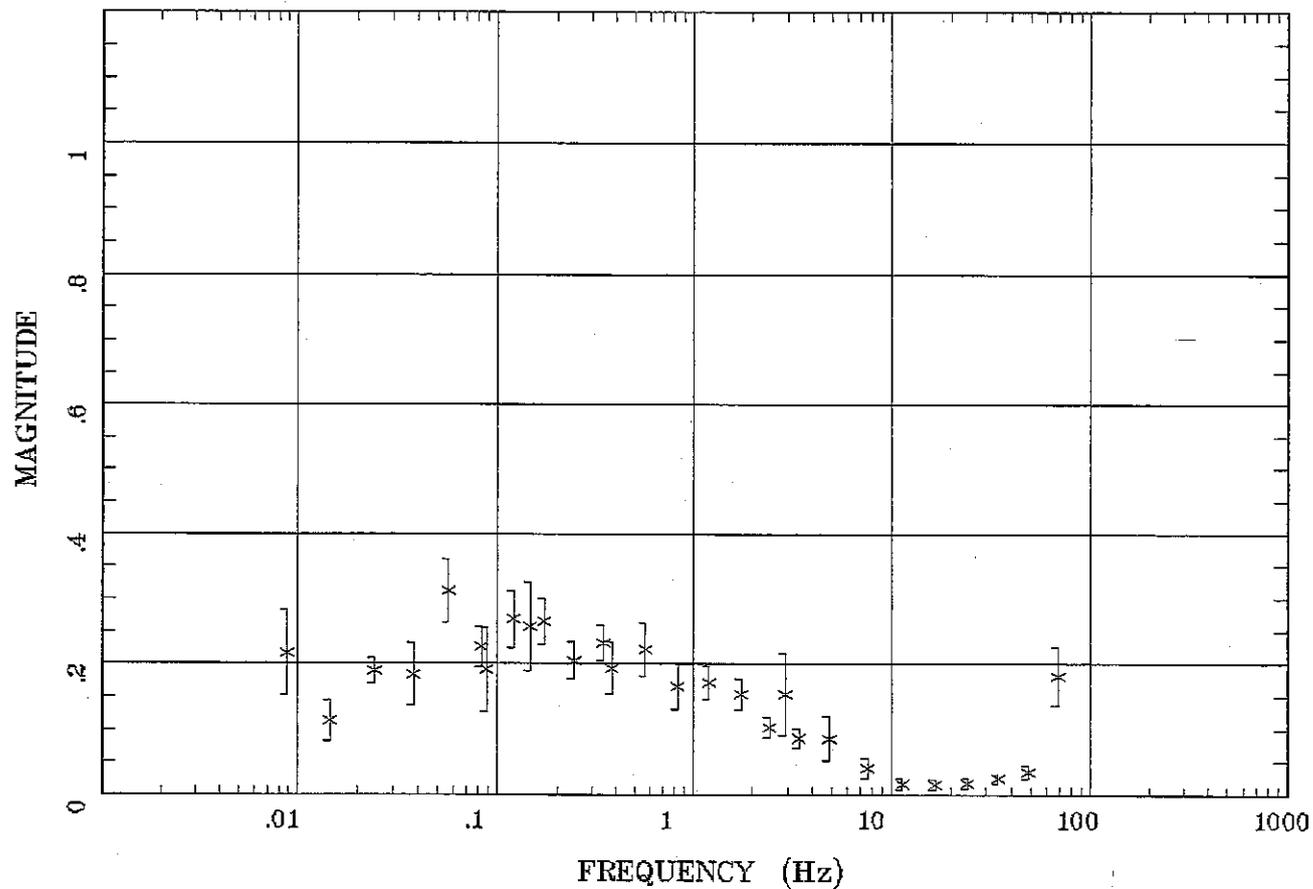
Client:
 Remote: none
 Acquired: 13:0 Jul 23, 2009
 Survey Co:USGS

Rotation:
 Filename: wp05m3.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:53 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 5

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



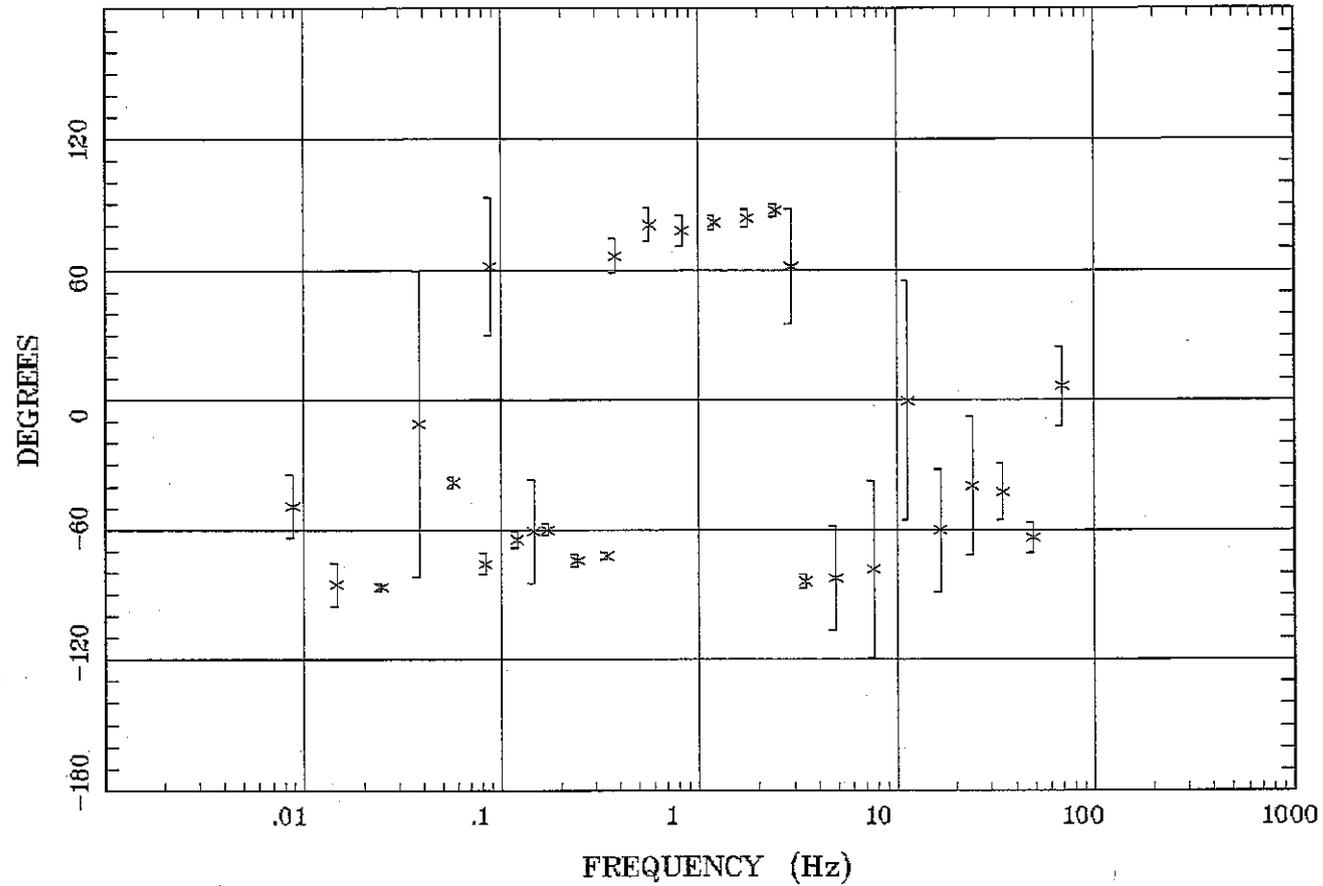
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

TIPPER STRIKE

Wheeler Peak Quad, NM



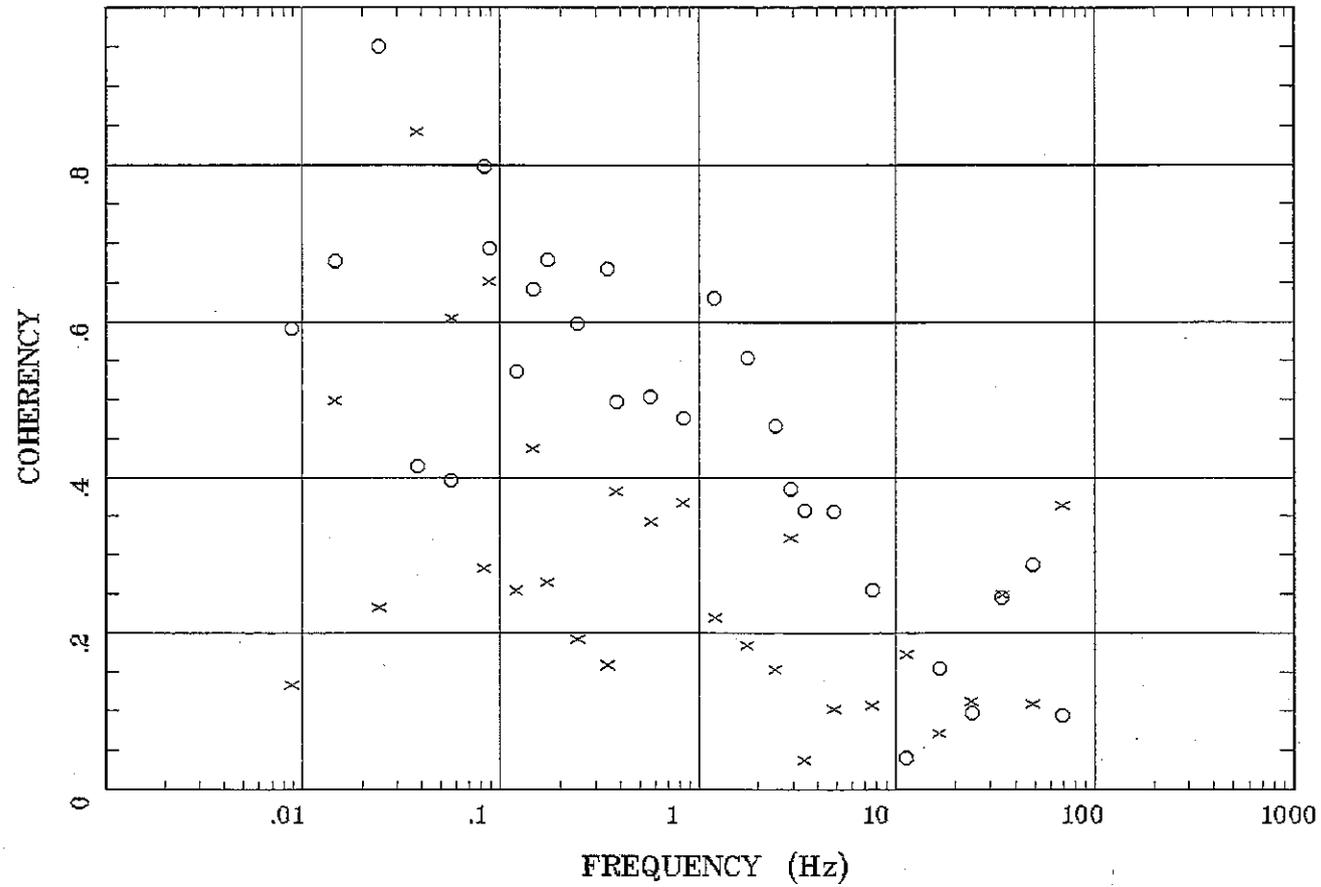
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 5

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



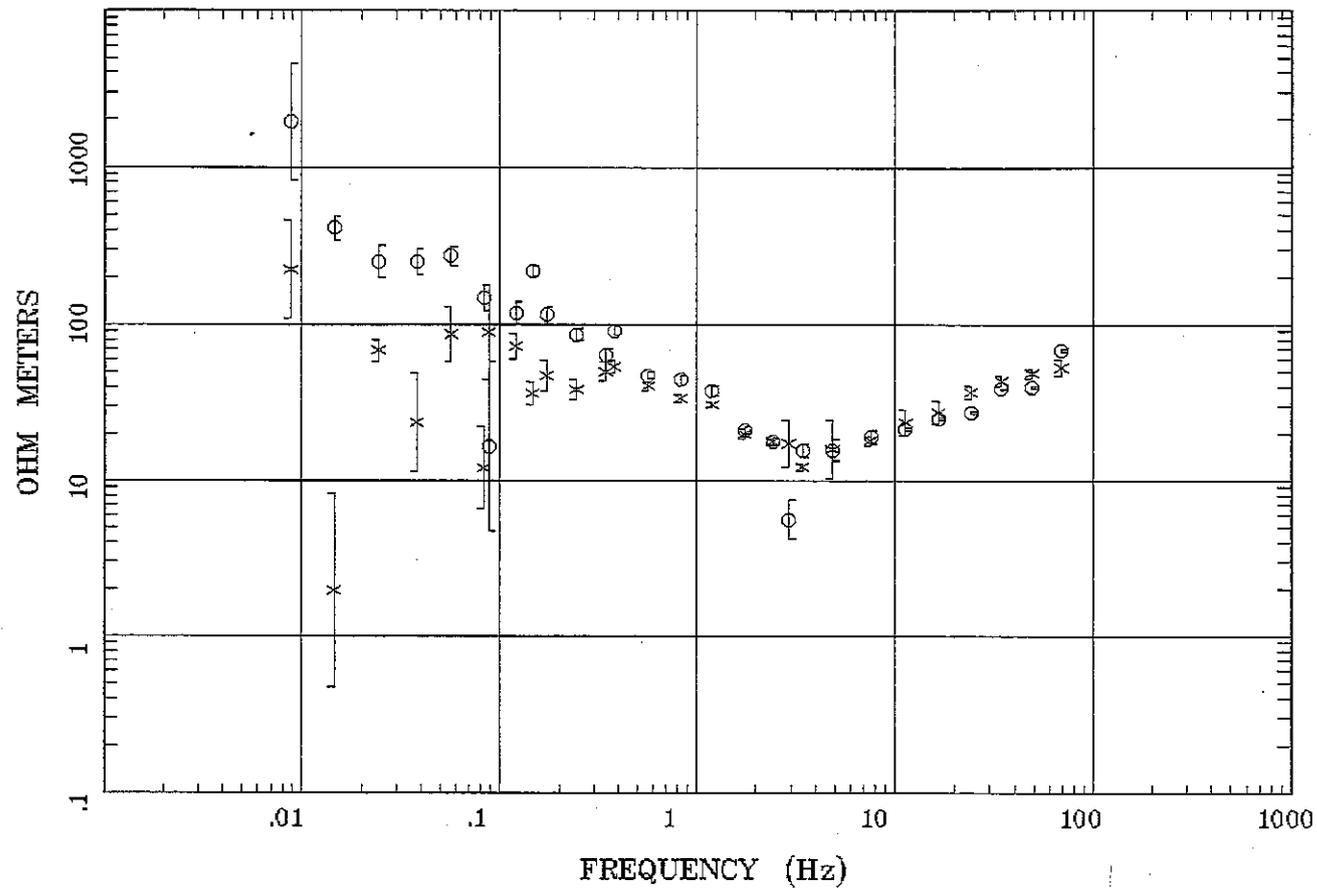
Client:
Remote: none
Acquired: 13:0 Jul 23, 2009
Survey Co:USGS

Rotation:
Filename: wp05m3.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:53 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 7

APPARENT RESISTIVITY

Wheeler Peak Quad, NM



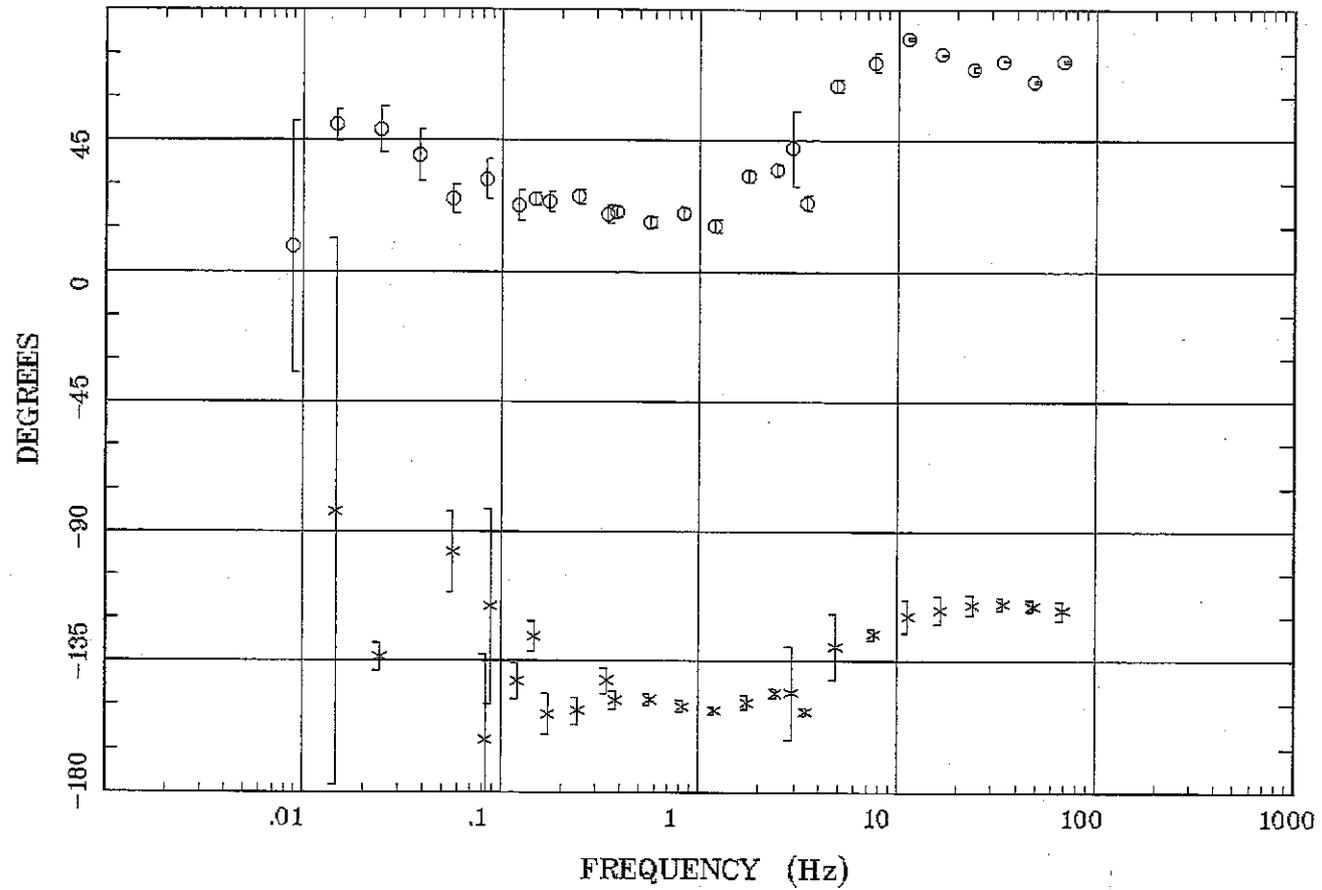
Client:
 Remote: none
 Acquired: 13:1 Jul 26, 2009
 Survey Co:USGS

Rotation:
 Filename: wp07m4.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:44 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 7

IMPEDANCE PHASE

Wheeler Peak Quad, NM



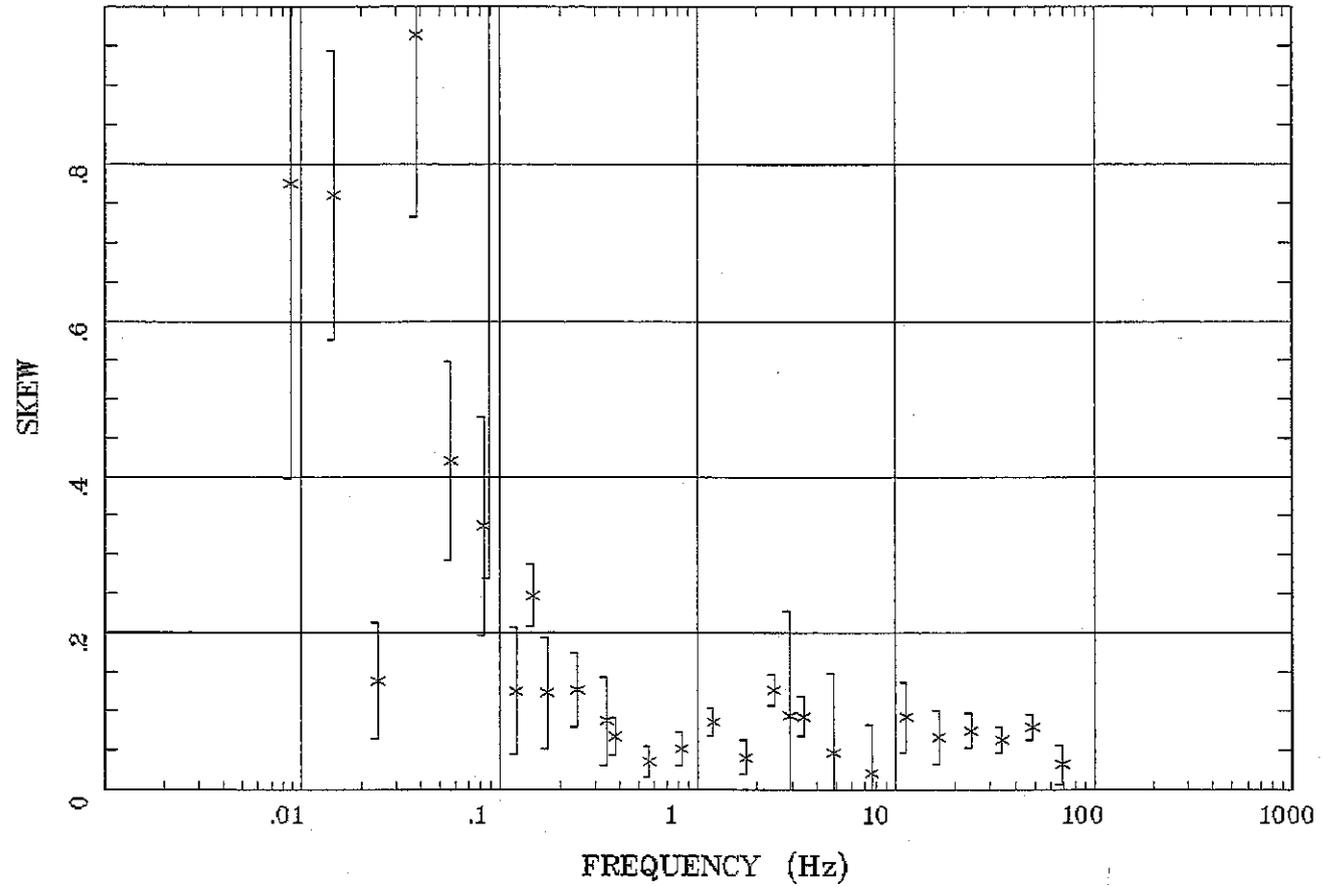
Client:
Remote: none
Acquired: 13:1 Jul 26, 2009
Survey Co:USGS

Rotation:
Filename: wp07m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:44 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 7

IMPEDANCE SKEW

Wheeler Peak Quad, NM



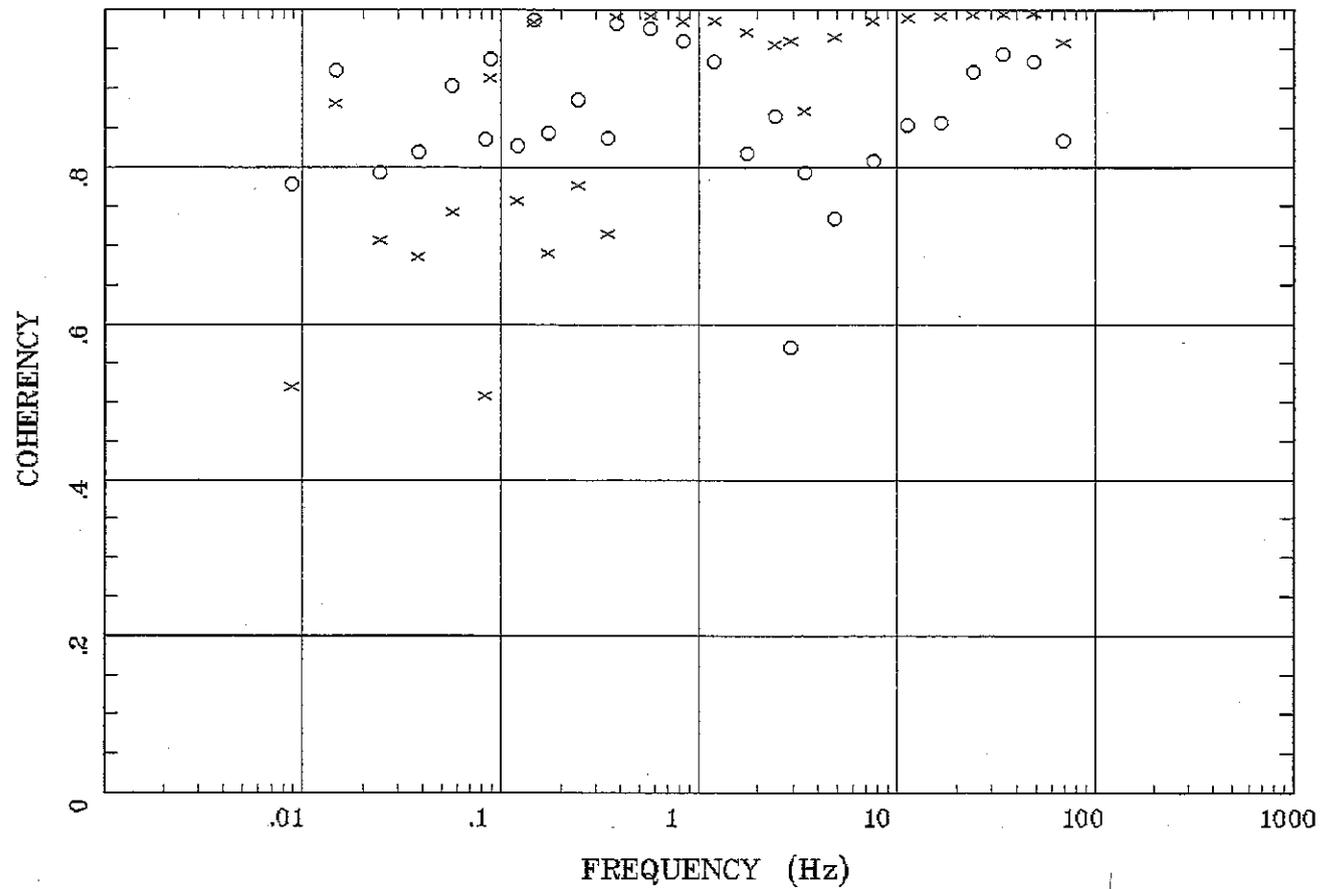
Client:
Remote: none
Acquired: 13:1 Jul 26, 2009
Survey Co:USGS

Rotation:
Filename: wp07m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:44 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 7

E MULT Coh.

Wheeler Peak Quad, NM



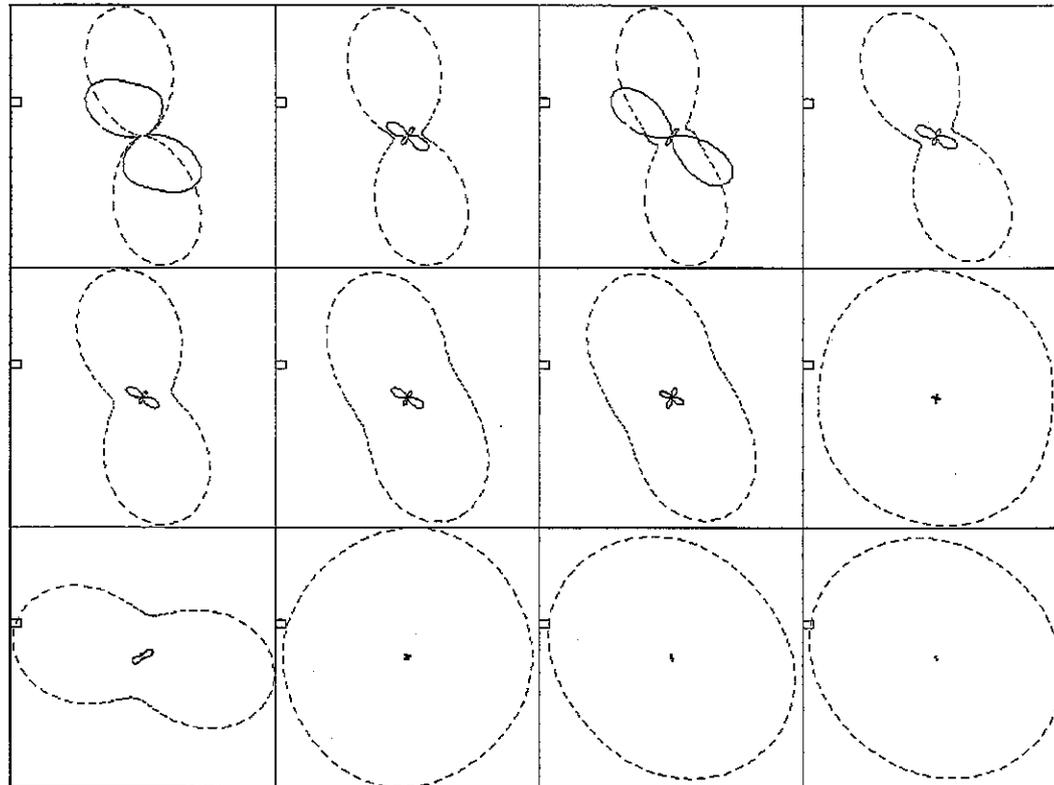
Client:
Remote: none
Acquired: 13:1 Jul 26, 2009
Survey Co:USGS

Rotation:
Filename: wp07m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:44 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 7

POLAR PLOTS

Wheeler Peak Quad, NM



.0088 Hz	.0244 Hz	.0566 Hz	.120 Hz
.172 Hz	.345 Hz	.566 Hz	1.758 Hz
2.930 Hz	7.617 Hz	16.602 Hz	34.375 Hz

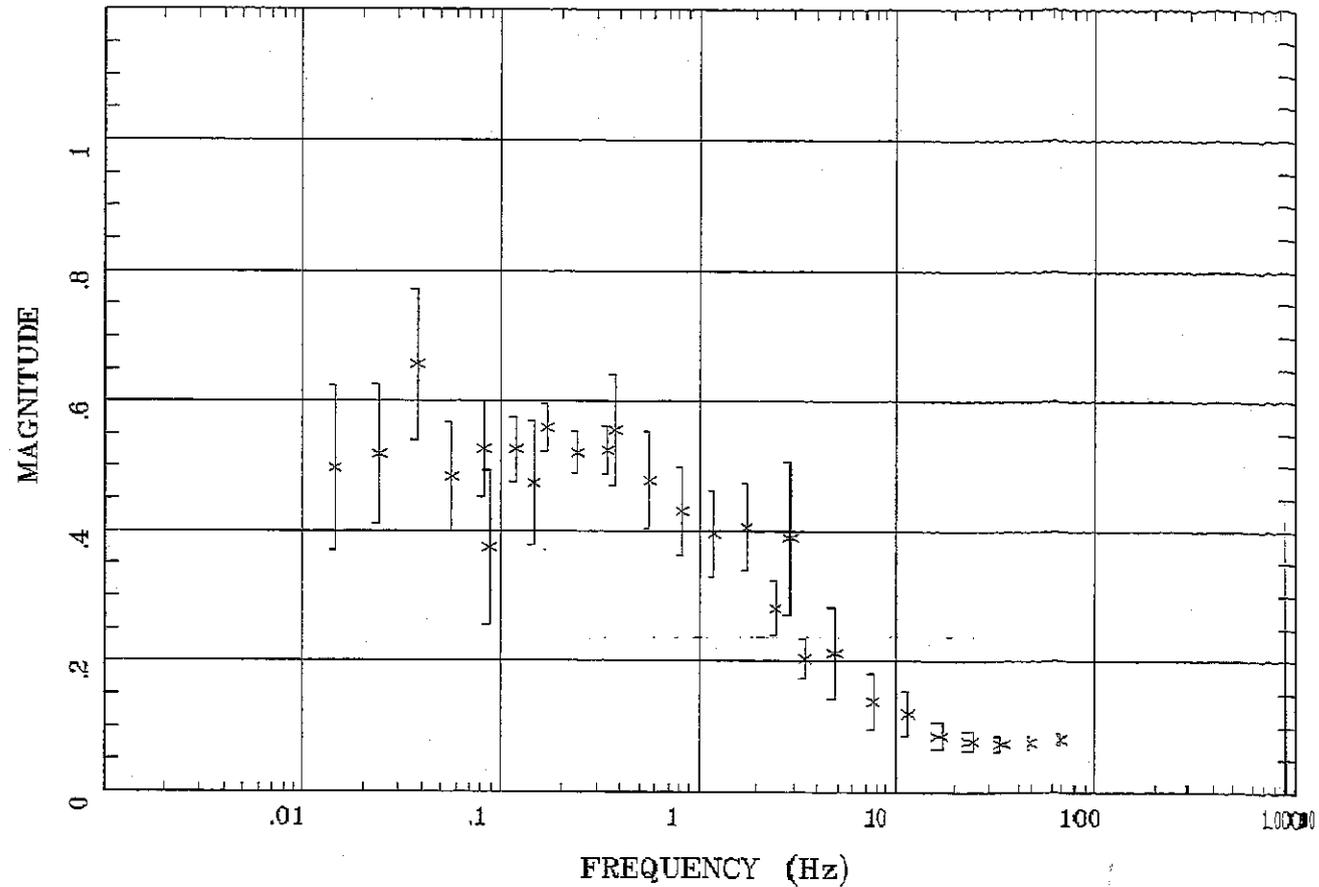
Client:
 Remote: none
 Acquired: 13:1 Jul 26, 2009
 Survey Co:USGS

Rotation:
 Filename: wp07m4.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:44 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 7

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



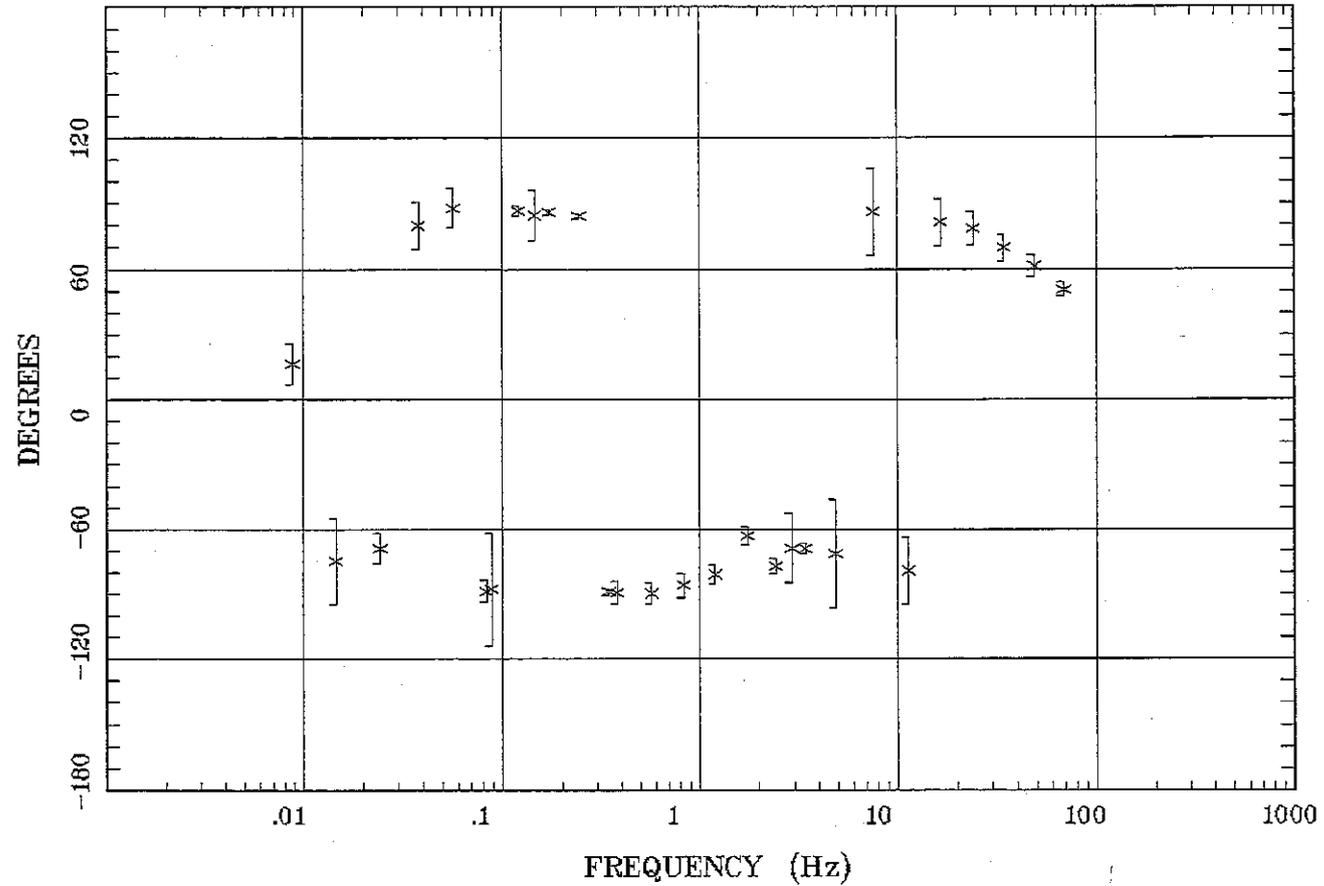
Client:
Remote: none
Acquired: 13:1 Jul 26, 2009
Survey Co:USGS

Rotation:
Filename: wp07m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch6 Ch7
Plotted: 12:44 Feb 12, 2010
< EMI - Electro Magnetic Instrumentss >

Station 7

TIPPER STRIKE

Wheeler Peak Quad, NM



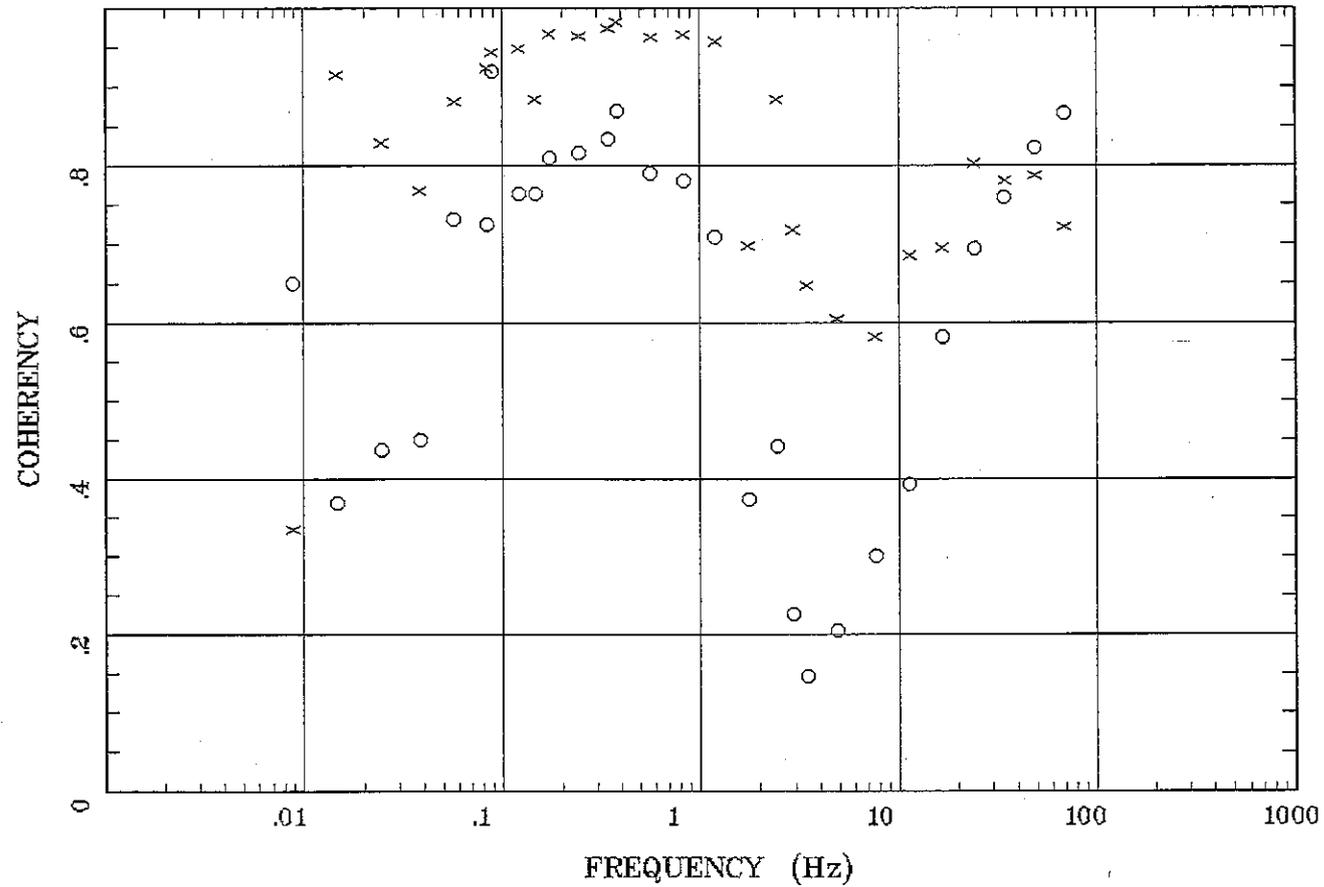
Client:
Remote: none
Acquired: 13:1 Jul 26, 2009
Survey Co:USGS

Rotation:
Filename: wp07m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:44 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 7

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



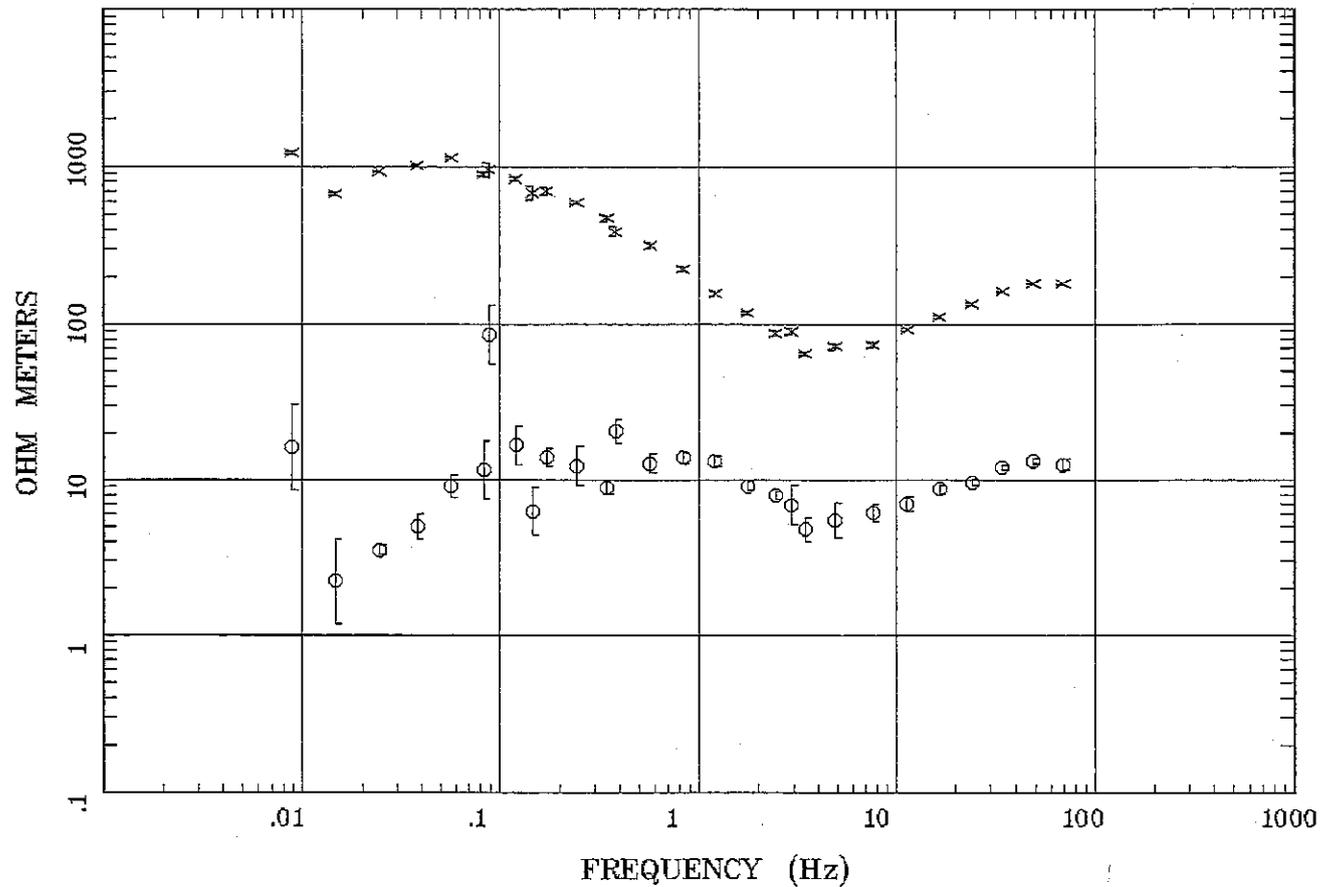
Client:
Remote: none
Acquired: 13:1 Jul 26, 2009
Survey Co:USGS

Rotation:
Filename: wp07m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:44 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

APPARENT RESISTIVITY

Wheeler Peak Quad, NM



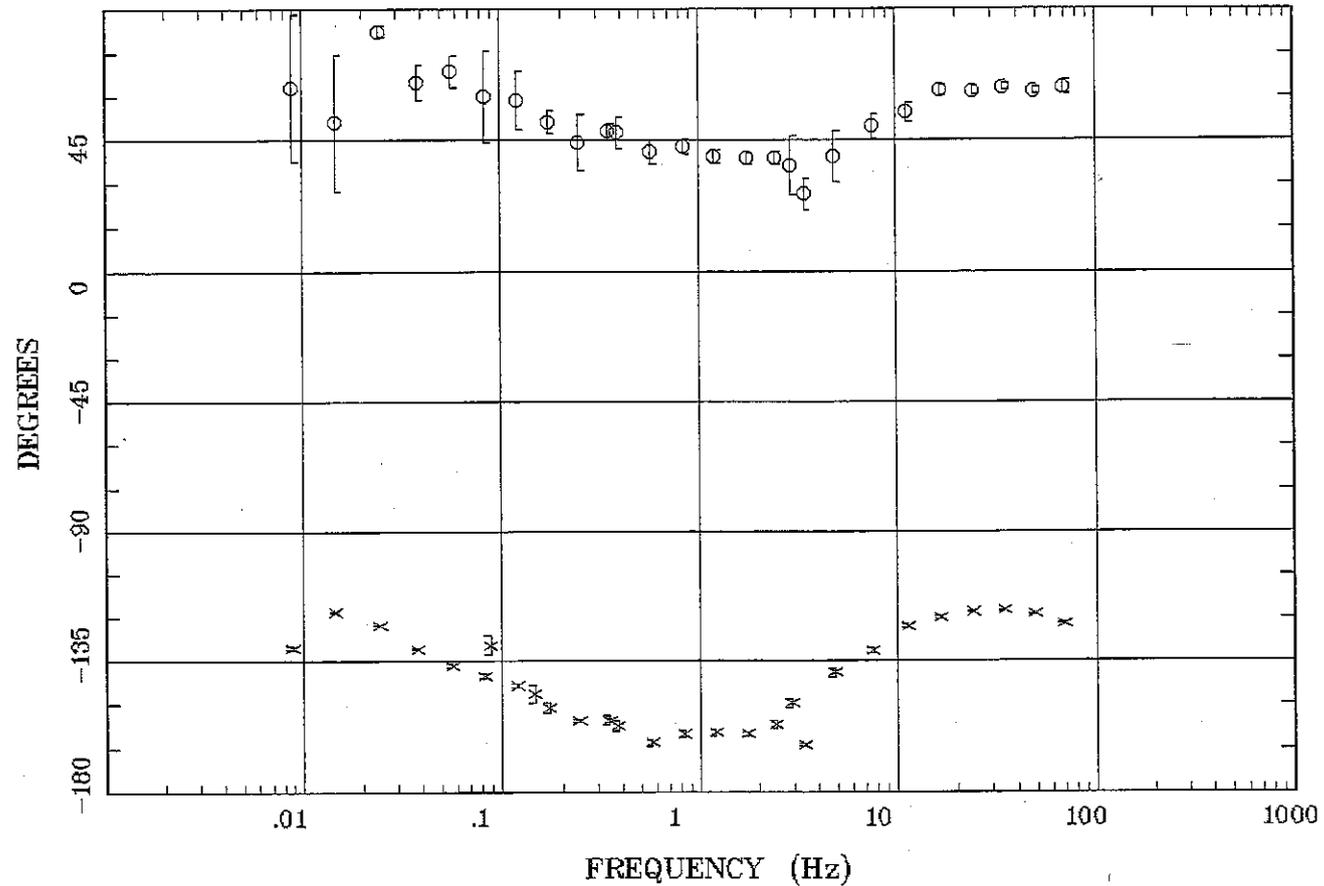
Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

IMPEDANCE PHASE

Wheeler Peak Quad, NM



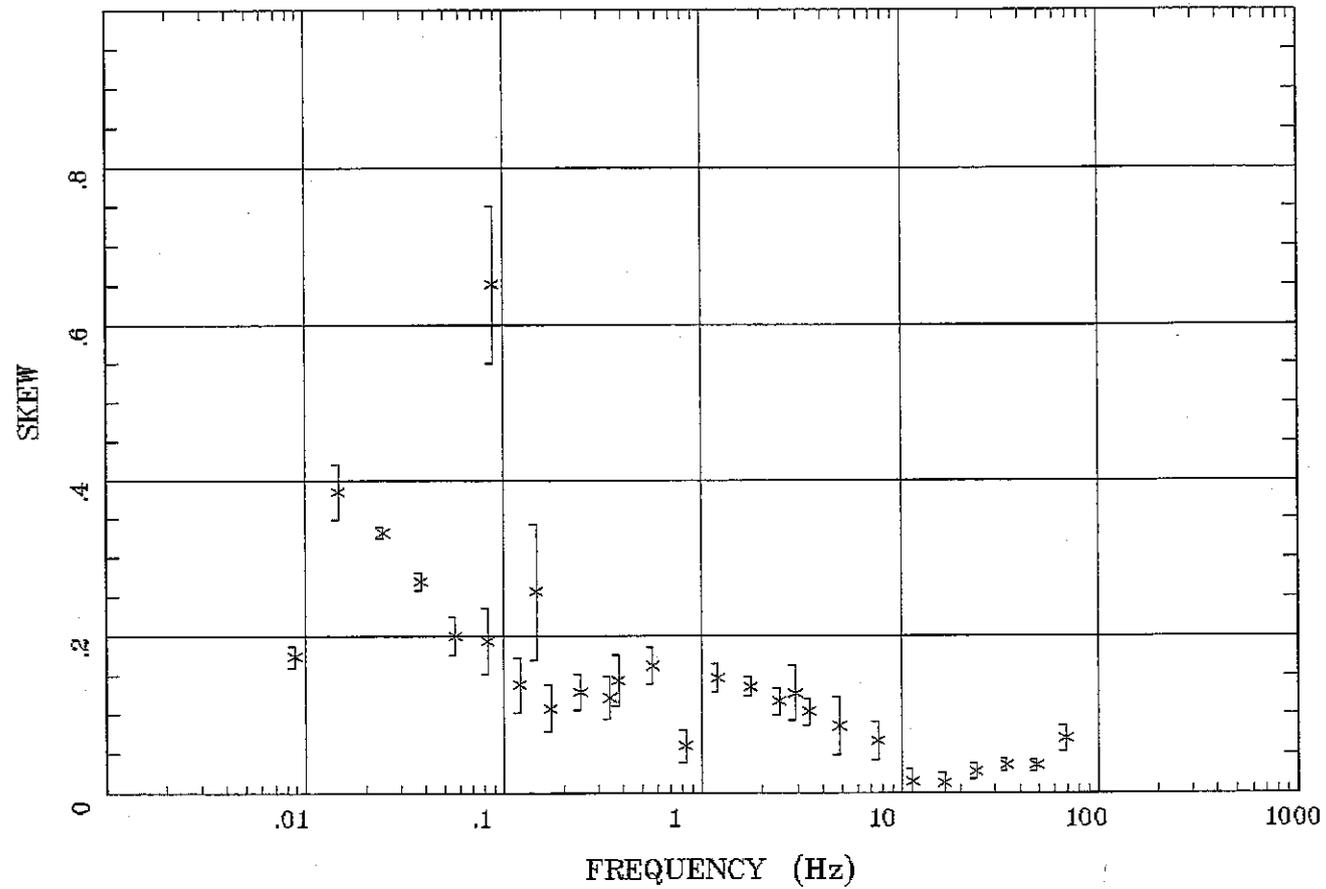
Client:
 Remote: none
 Acquired: 15:5 Jul 25, 2009
 Survey Co:USGS

Rotation:
 Filename: wp09m4.avg
 Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
 Plotted: 12:54 Feb 12, 2010
 < EMI - ElectroMagnetic Instruments >

Station 9

IMPEDANCE SKEW

Wheeler Peak Quad, NM



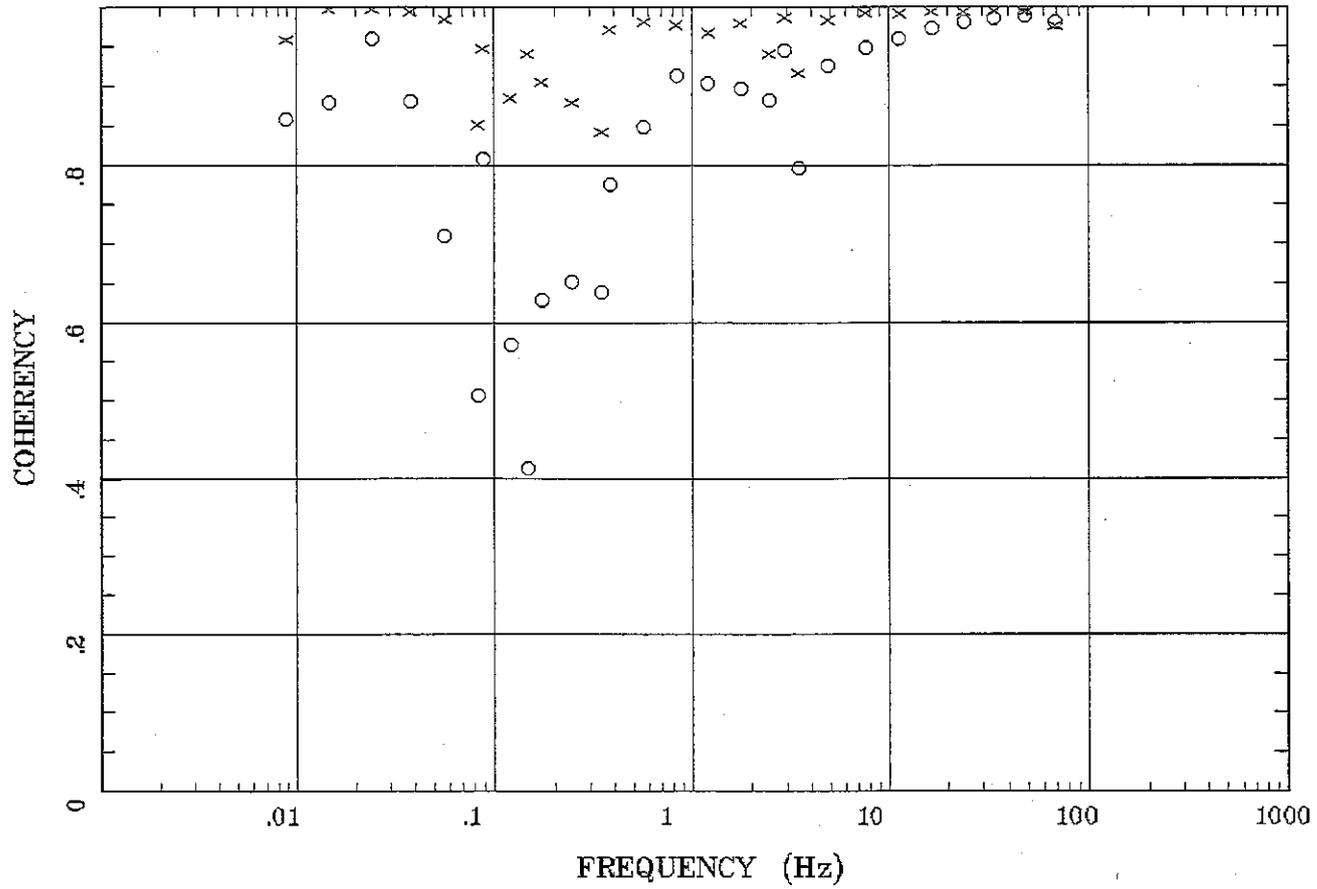
Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

E MULT Coh.

Wheeler Peak Quad, NM



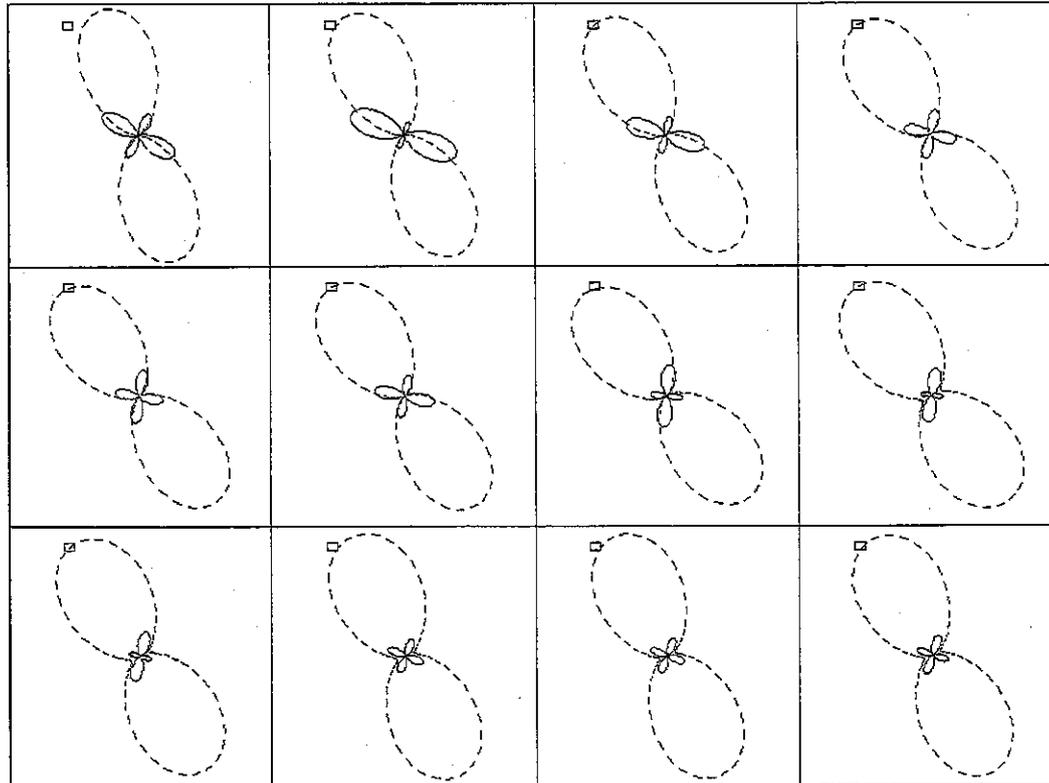
Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

POLAR PLOTS

Wheeler Peak Quad, NM



.0088 Hz
.172 Hz
2.930 Hz

.0244 Hz
.345 Hz
7.617 Hz

.0566 Hz
.566 Hz
16.602 Hz

.120 Hz
1.758 Hz
34.375 Hz

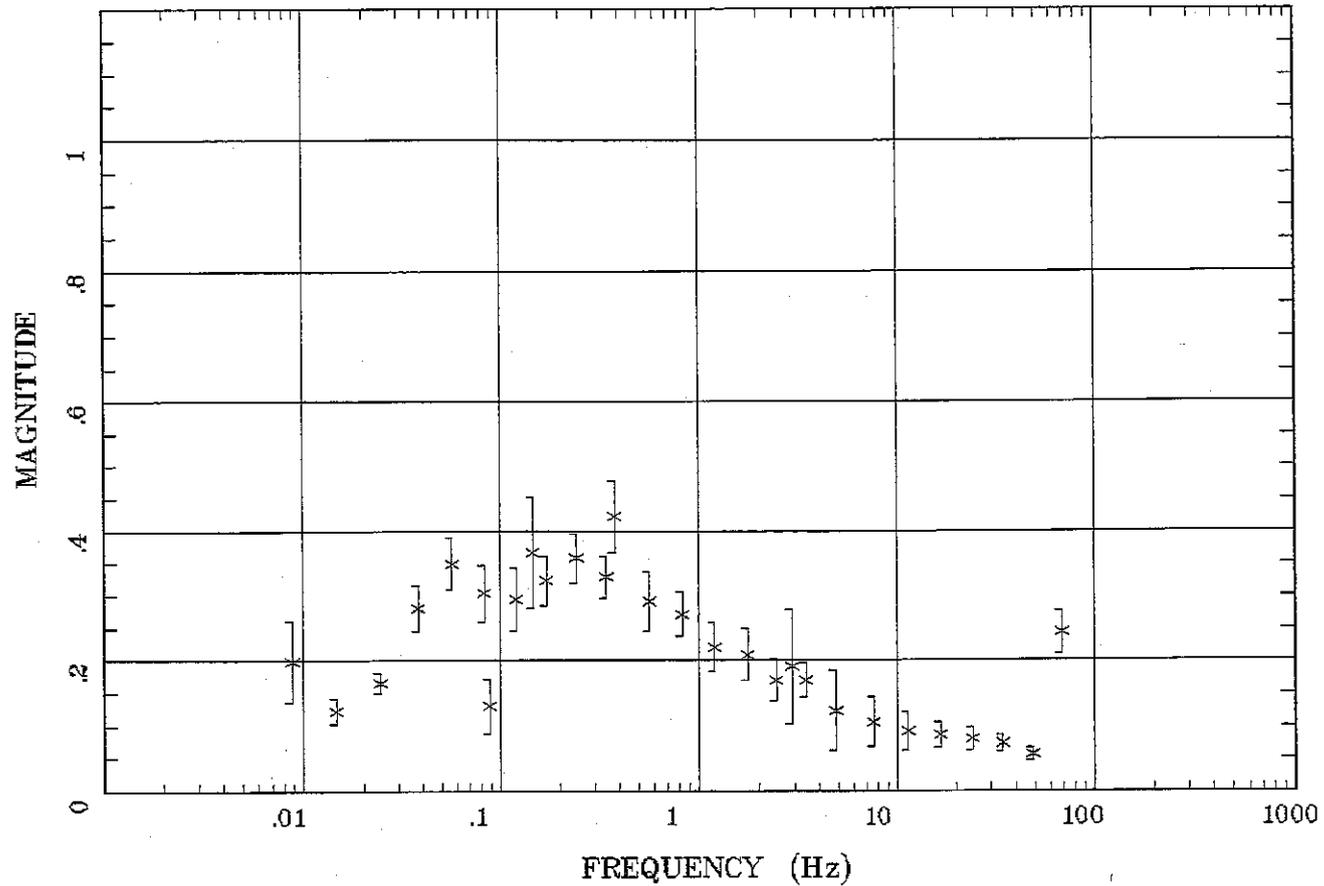
Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

TIPPER MAGNITUDE

Wheeler Peak Quad, NM



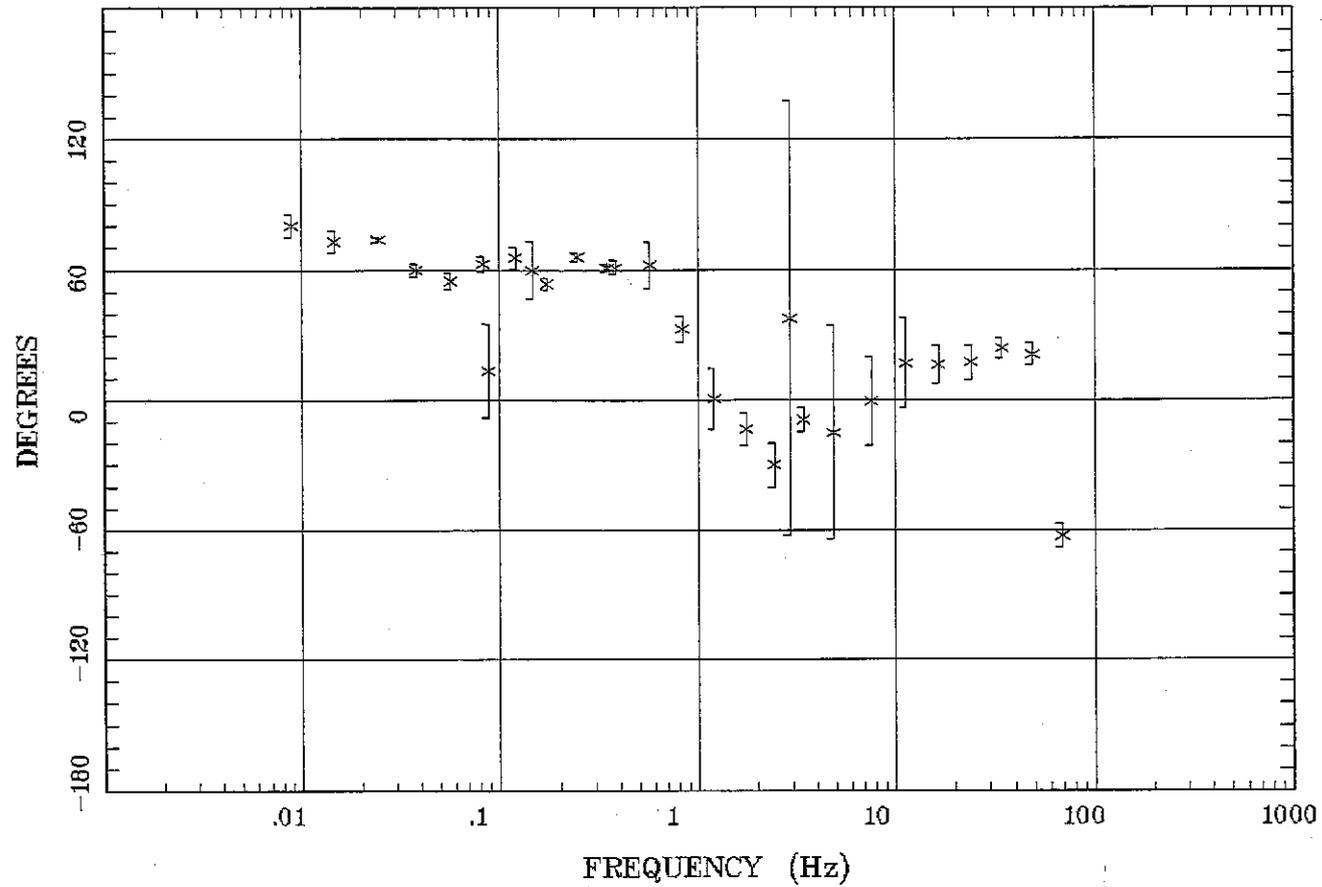
Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

TIPPER STRIKE

Wheeler Peak Quad, NM



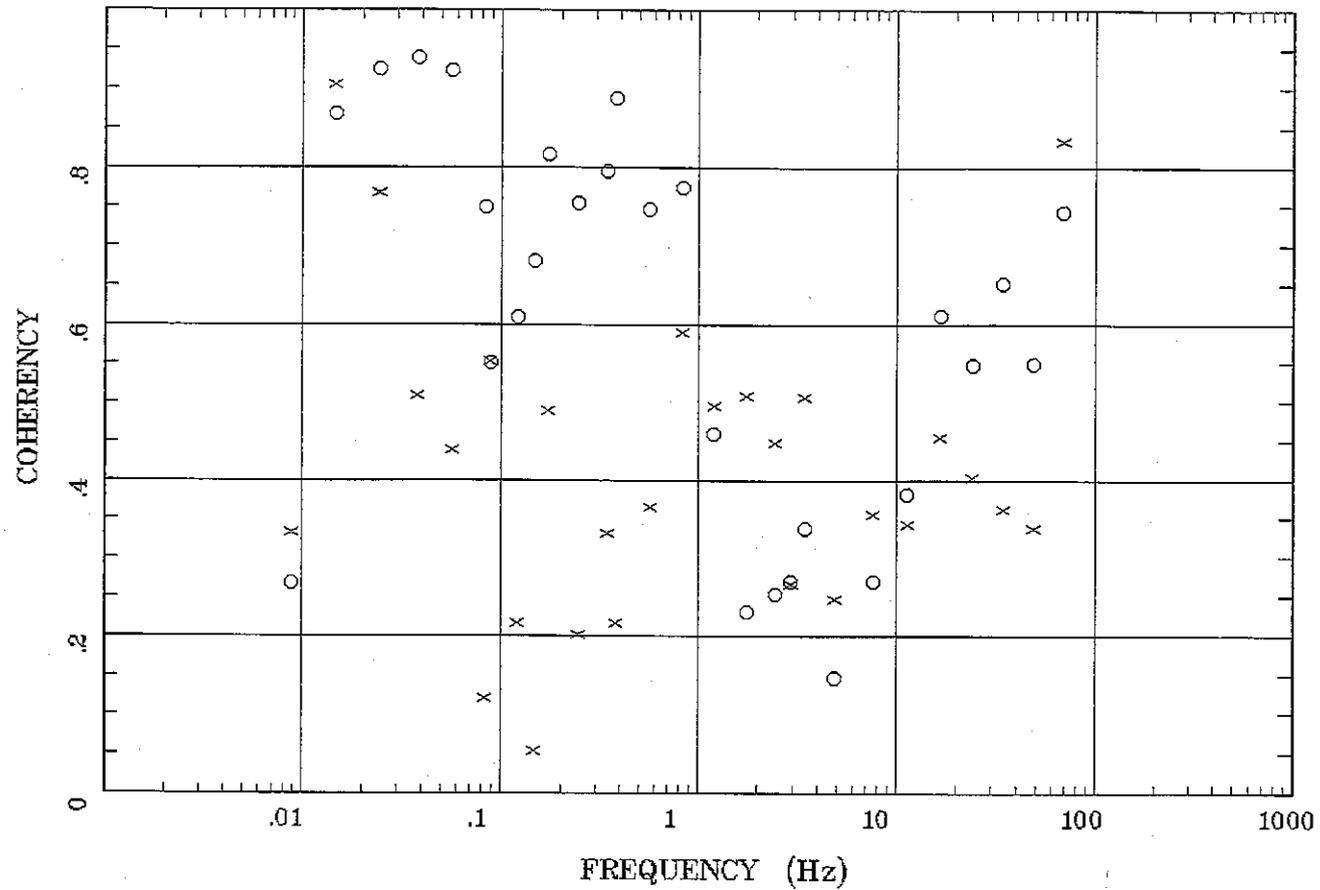
Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
< EMI - ElectroMagnetic Instruments >

Station 9

HzHx.x Coh HzHy.o

Wheeler Peak Quad, NM



Client:
Remote: none
Acquired: 15:5 Jul 25, 2009
Survey Co:USGS

Rotation:
Filename: wp09m4.avg
Channels: Ch1 Ch2 Ch3 Ch4 Ch5 Ch3 Ch4
Plotted: 12:54 Feb 12, 2010
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