

DISCUSSION

This sheet includes maps that show the interpreted thickness and depth to base of uppermost Pleistocene and Holocene deposits in California's State Waters for the offshore of Pigeon Point map area (Maps A, B), as well as for a larger area that extends along the coast from the Pigeon Point area (Maps A, B, C, D) to establish regional context. High-resolution seismic-reflection profiles (Fig. 1; see also, sheet 8) show a lower unit of deformed Neogene bedrock and one or two upper units that consist of upper Quaternary sediments. The bedrock-sediment contact is an angular unconformity that commonly is marked by minor channeling, an eastward-sloping onto reflection-free bedrock, and an upward change to lower amplitude, more diffuse reflections.

Two upper Quaternary units are recognized on high-resolution seismic-reflection profiles in the Pigeon Point to southern Monterey Bay region. Although not present in this map area, the lower unit notably includes low-amplitude, low-angle (1° to 3°), offshore-dipping clinoforms (Cattaneo, 2006). The upper unit (blue shading on Fig. 1; see also, Figs. 1, 2, 3, 5, 6, 8, 9 on sheet 8) typically is characterized by low-amplitude, continuous to moderately continuous, diffuse, subparallel reflections, and it has a maximum thickness of about 22 m. Our preferred hypothesis is that the lower of the two upper Quaternary sedimentary units represents a prograding delta and (or) shoreface complex that formed between about 30,000 and 21,000 years ago during the pre-LGM (Last Glacial Maximum) sea-level drop of marine-isotope stage 2 (Waelbroeck and others, 2002). The overlying upper unit (blue shading) represents shelf deposits that formed during the post-LGM sea-level rise of the last about 21,000 years (Stanford and others, 2011). In this interpretation, the surface at the top of the lower, clinoform-bearing unit is a transgressive surface of erosion that formed as the shoreface migrated landward. Because these two upper Quaternary units each consist of unconformably overlain Quaternary sediments and together overlie the prominent angular unconformity with bedrock, we have combined their thicknesses on Maps B and D.

To make these maps, water bottom and depth to base of the uppermost Pleistocene and Holocene sediment layer were mapped from seismic-reflection profiles (Fig. 1; see also, sheet 8). The difference between the two horizons was exposed for every shot point as XY coordinates (UTM zone 10) and two-way travel time (TWT). The thickness of the post-LGM unit (Map B, D) was determined by applying a sound velocity of 1,600 m/sec to the TWT. The thickness points were interpolated to a preliminary continuous surface, overlain with zero-thickness bedrock outcrops (see sheet 10), and contoured, following the methodology of Wong and others (2012).

The thickness of the post-LGM sediments in the offshore of Pigeon Point map area ranges from 0 to 22 m (Map B), and the depth to the unconformity at the base of this unit ranges from less than 10 to 75 m (Map A). Mean sediment thickness for the map area is 6 m, and the total sediment volume is 765 × 10⁹ m³ (see table 7-1 in pamphlet). The thickest sediment (22 m) is found in the northwest corner of the map area (Map B), coincident with a decrease in slope on the bedrock unconformity (Maps A, C). Sediment thickness diminishes to the southeast, and sediment cover offshore of Point Año Nuevo is either thin or nonexistent (Map B, D). The extensive bedrock outcrop in this area coincides with an area of transposition between the Accension Fault and the San Gregorio Fault Zone. The relative lack of offshore sediment probably is a result of both low sediment supply and ongoing exposure to erosive wave energy.

Six different informal "domains" of thickness of uppermost Pleistocene to Holocene sediment (table 7-1) are recognized on the regional sediment thickness map (Map D on sheet 9), each with its own diverse set of geologic and (or) oceanographic controls. Note that data from within the Monterey Canyon system (including Sausal Canyon), in the southern part of the Pigeon Point to southern Monterey Bay region, were excluded from this analysis because available seismic-reflection data are insufficient to map sediment distribution in this extremely variable environment.

(1) The southern Monterey Bay domain is bounded by the Monterey Bay shoreline on the south and east, the Monterey Canyon on the north, and the limit of California's State Waters on the west. Sediment derived from the Salinas River forms a large, shore-parallel, subaqueous delta (discussed as much as 32 m) that progrades across a thinly sediment-mantled bedrock shelf. Small changes in sediment thickness on the shelf are controlled by irregular bedrock relief that at least partly attributable to the Monterey Bay Fault Zone (Greene, 1990).

(2) The northern Monterey Bay domain is bounded on the south by Monterey Canyon, on the north and east by the Monterey Bay shoreline, and on the west by the limit of California's State Waters. The head of Monterey Canyon extends nearly to the shoreline, and the canyon forms a sediment trap that effectively separates the linear- and shelf-sediment transport systems of the northern and southern Monterey Bay domains. The northern Monterey Bay domain is characterized by (a) a sediment-poor inner shelf cut by paleochannels of the San Lorenzo River, the Pajaro River, and Sequoia Creek; (b) a midshelf depositor that has sediment as thick as 32 m, much of which was deposited in a pre-LGM prograding delta and (or) shoreface complex and was preserved above a decrease in slope on the underlying unconformity; and (c) a midshelf to outer shelf zone in which sediment generally becomes progressively thinner in the offshore direction.

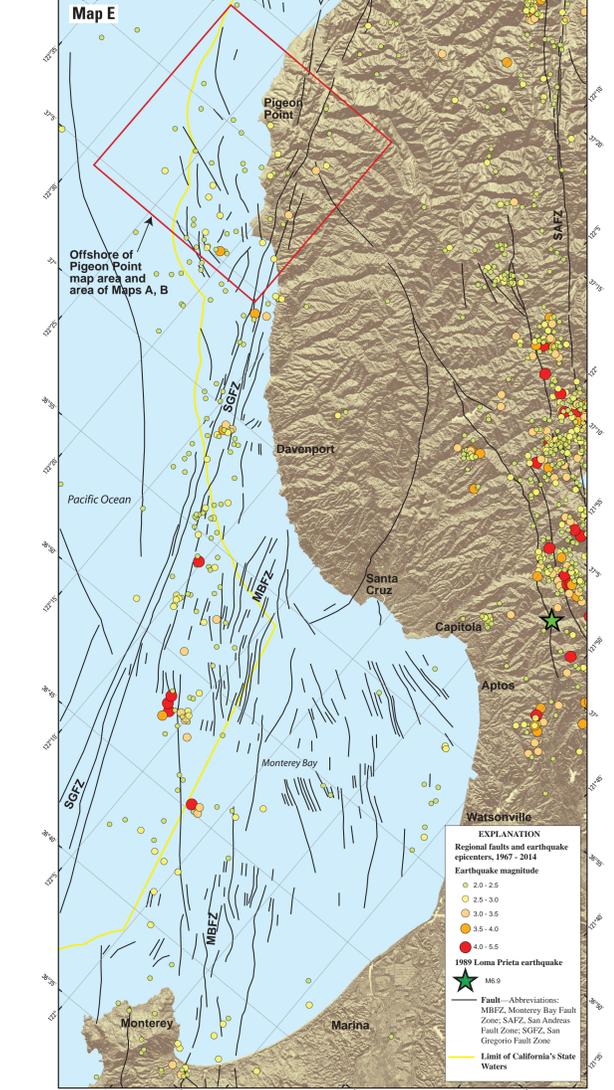
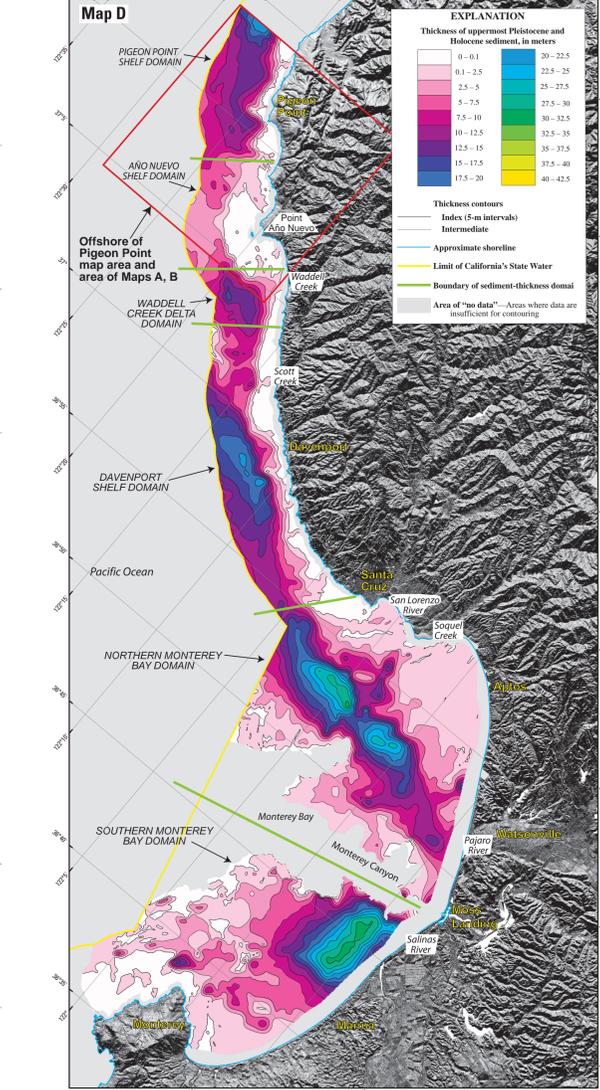
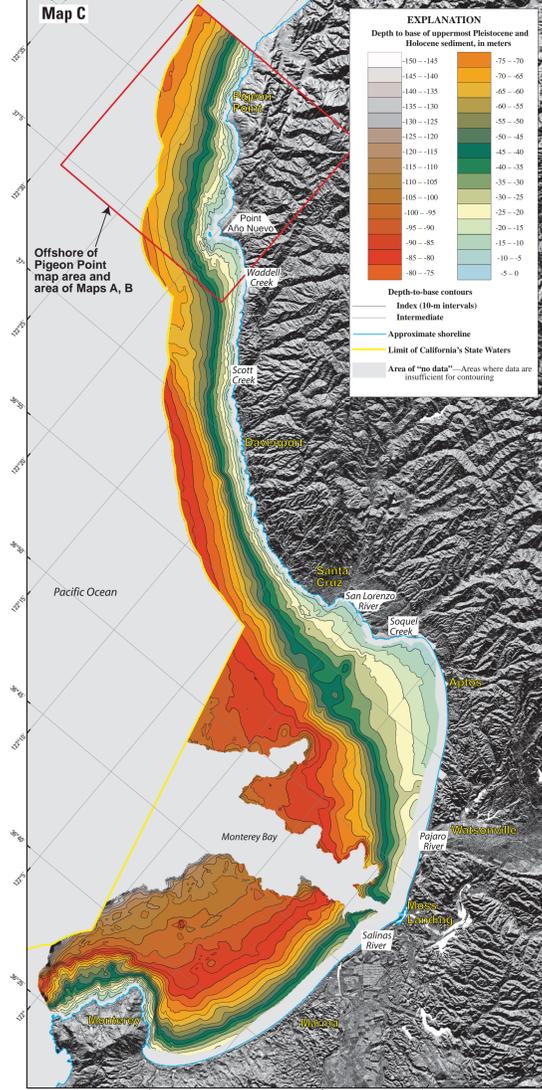
(3) The Davenport shelf domain extends from the northern limit of Monterey Bay northward to the southern margin of the Waddell Creek depositor. The Davenport shelf domain, as well as the three domains farther north, occupy a section of open, wave-dominated coast that is exposed to wave energy higher than that of the Monterey Bay domains to the south. The Davenport shelf domain includes the Davenport depositor, a prominent midshelf, shore-parallel depositor present between Davenport and Santa Cruz that mostly consists of a lower, pre-LGM, clinoform-bearing unit of inferred prograding-shoreface origin. Sediment in this depositor also is preserved in accommodation space linked to an area of increase in the slope of the underlying unconformity. Sediment thickness within the Davenport shelf domain decreases to both the northwest and southeast of this depositor, owing to the presence of elevated bedrock and (or) the relative absence of the lower clinoform-bearing unit.

(4) The Waddell Creek delta domain lies offshore of the mouth of the Waddell Creek coastal watershed, and it is connected to it by a submerged channel. The domain is both distinguished and delineated by the significant Waddell Creek depositor (maximum sediment thickness of 19 m), which forms a moundlike delta that consists entirely of inferred post-LGM deposits whose primary source is Waddell Creek. Sediment thin both north and south of this moundlike delta; its preservation is attributable to its emplacement from erosive wave energy located on the south flank of Point Año Nuevo.

(5) The Año Nuevo shelf domain lies offshore of Point Año Nuevo, from just north of Franklin Point on the north to just north of the mouth of Waddell Creek on the south. Bedrock exposures, which locally reach water depths of 45 m, cover a substantial part of this wave-exposed domain; in deeper waters farther offshore, sediment thickness in this domain appears to be limited both by the lack of sediment supply (because of its distance from large coastal watersheds) and by the presence of uplifted bedrock, which is linked to a local zone of transposition in the San Gregorio Fault Zone (Weber, 1990). The uplift has raised this domain and exposed it to the high wave energy that is characteristic of this area (Storazzi and Wingfield, 2005).

(6) The Pigeon Point shelf domain lies on the west flank of the Pigeon Point high (McCulloch, 1987). Sediment in the Pigeon Point shelf domain is thickest in a shore-parallel band that overlies a slope break in the underlying bedrock surface. Much of the sediment probably was derived from Pescadero Creek, a large coastal watershed that enters the Pacific Ocean about 3 km north of the Pigeon Point to southern Monterey Bay regional map area (see Map C, D on sheet 9). The Pigeon Point shelf domain is transitional to the Pacific-Pescadero shelf domain just north of it (see Watt and others, 2014).

Map E shows the regional pattern of major faults and earthquakes occurring between 1967 and April 2014 that have inferred or measured magnitudes of 2.0 and greater. Fault locations, which have been simplified, are compiled from our mapping within California's State Waters (see sheet 10), from Wagner and others (2002), and from the U.S. Geological Survey's Quaternary fault and fold database (U.S. Geological Survey and California Geological Survey, 2010). Earthquake epicenters are from the Northern California Earthquake Data Center (2014), which is maintained by the U.S. Geological Survey and the University of California, Berkeley, Seismological Laboratory. The 1909 Loma Prieta earthquake (M_{6.9}, 10/17/1989), on the San Andreas Fault Zone in the Santa Cruz Mountains (Spudis, 1990), is the most significant event in the region. In the offshore of Pigeon Point map area, a substantial amount of seismicity is broadly distributed between the San Gregorio Fault Zone and the Accension Fault—the largest recorded earthquake (M_{5.6}) having occurred on 6/26/1991.



Map A: Depth to base of uppermost Pleistocene and Holocene sediment, in meters. Map B: Thickness of uppermost Pleistocene and Holocene sediment, in meters. Map C: Offshore of Pigeon Point map area and area of Maps A, B. Map D: Offshore of Pigeon Point map area and area of Maps A, B. Map E: Offshore of Pigeon Point map area and area of Maps A, B.

Map C: Offshore of Pigeon Point map area and area of Maps A, B. Map D: Offshore of Pigeon Point map area and area of Maps A, B. Map E: Offshore of Pigeon Point map area and area of Maps A, B.

Map D: Offshore of Pigeon Point map area and area of Maps A, B. Map E: Offshore of Pigeon Point map area and area of Maps A, B.

Map E: Offshore of Pigeon Point map area and area of Maps A, B.

Local (Offshore of Pigeon Point Map Area) and Regional (Offshore from Pigeon Point to Southern Monterey Bay) Shallow-Subsurface Geology and Structure, California

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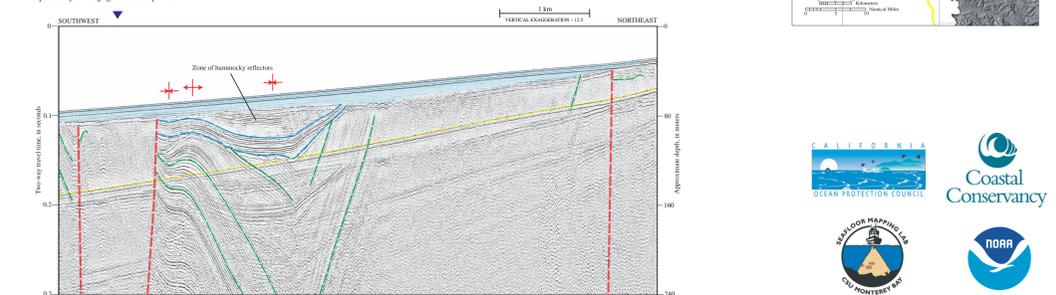


Figure 1: USGS high-resolution minisepikar seismic-reflection profile P19-25 collected in 2010 on survey S-15-10-NC, see fig. 5 on sheet 8, which crosses shelf south of Pigeon Point (see Map A for location). Profile highlights faulted and folded strata beneath continental shelf. Dashed red lines show faults. Magenta symbols show fold axes (overlying arrows, anticlines; overlying arrows, synclines). Blue shading shows inferred uppermost Pleistocene and Holocene strata, deposited since last sea-level lowered about 21,000 years ago. Underlying reflectors image deformed Tertiary and Cretaceous strata. Dashed blue lines show unconformities. Note zone of hummocky reflectors between offshore parallel to subparallel reflectors near middle of profile (see cross-line, Fig. 3 on sheet 8). Dashed green lines highlight some continuous reflections that reveal structure (note distinctive stratigraphic markers). Dashed yellow lines is seafloor multiple (echo of seafloor reflector). Purple triangle shows location of California's State Waters limit (yellow line on Maps A, B).



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