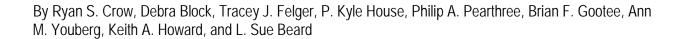


The Colorado River and its Deposits Downstream from Grand Canyon in Arizona, California, and Nevada



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The Colorado River and Its Deposits Downstream from Grand Canyon in Arizona, California, and Nevada

By Ryan S. Crow¹, Debra Block¹, Tracey J. Felger¹, P. Kyle House¹, Philip A. Pearthree², Brian F.Gootee², Ann M. Youberg², Keith A. Howard¹, and L. Sue Beard¹

Abstract

Understanding the evolution of the Colorado River system has direct implications for (1) the processes and timing of continental-scale river system integration, (2) the formation of iconic landscapes like those in and around Grand Canyon, and (3) the availability of groundwater resources. Spatial patterns in the position and type of Colorado River deposits, only discernible through geologic mapping, can be used to test models related to Colorado River evolution. This is particularly true downstream from Grand Canyon where ancestral Colorado River deposits are well-exposed. We are principally interested in (1) regional patterns in the minimum and maximum elevation of each depositional unit, which are affected by depositional mechanism and postdepositional deformation; and (2) the volume of each unit, which reflects regional changes in erosion, transport efficiency, and accommodation space. The volume of Colorado River deposits below Grand Canyon has implications for groundwater resources, as the primary regional aquifer there is composed of those deposits. To this end, we are presently mapping Colorado River deposits and compiling and updating older mapping. This preliminary data release shows the current status of our mapping and compilation efforts. We plan to update it at regular intervals in conjunction with ongoing mapping.

Background

Ongoing geologic mapping campaigns in the lower Colorado River corridor (LOCO), led by the U.S. Geological Survey (USGS) and the Arizona Geological Survey (AZGS), are focused on the identification and characterization of Colorado River deposits in the corridor between Hoover and Imperial Dams, in Nevada, California, and Arizona (fig. 1). Although these deposits have been studied for nearly 150 years (for example, Ives, 1861; Lee, 1908; Blackwelder, 1934; Longwell, 1936), there is continuing interest in the processes of river integration (for example, House and others, 2008; McDougall and Martinez, 2014; Spencer and others, 2013), the effect of tectonics and climate on Colorado River evolution (for example, Lucchitta, 1979; Malmon and others, 2011; Howard and others, 2015; Thacker and others, 2017), the timing of landscape evolution (for example, Howard and others, 2015; Kimbrough and others, 2015), and the extent of the aquifers housed in Colorado River deposits (for example, Owen-Joyce and others, 2000; Wiele and others, 2008). These ongoing research topics are greatly informed by detailed outcrop-scale studies, analytical studies, and an understanding of the regional patterns in the extent of

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Colorado River deposits. Field and subsurface observations (for example, Metzger, 1968; Buising, 1990; House and others, 2005; House and others, 2008; Pearthree and House, 2014; Gootee and others, 2016), and geochronology (for example, House and others, 2008; Dorsey and others, 2011; Sarna-Wojcicki and others, 2011) have led to an increasingly refined understanding of the sequence and timing of river incision and deposition in the area. The main stratigraphic packages deposited by the Colorado River have been grouped by House (2016) as the LOCO strata, which includes (from oldest to youngest): the Bouse Formation, the Bullhead Alluvium, the Palo Verde alluvium, the Chemehuevi Formation, the Riverside alluvium (not yet included in our LOCO group compilation), and the Blythe alluvium. This new stratigraphic framework builds on the seminal work of Metzger (Metzger, 1968; Metzger and others, 1973; Metzger and Loeltz, 1973).

The Bouse Formation consists of carbonates and interbedded siliciclastic mud and sand, including abundant Colorado Plateau derived grains that were deposited as the Colorado River filled and spilled southward through a series of formerly internally drained basins to the proto-Gulf of California (Metzger, 1968; Buising, 1990; House and others, 2008; Pearthree and House, 2014; Spencer and others, 2013; Kimbrough and others, 2015). Previous work indicates a maximum age for the Bouse Formation in Cottonwood and Mohave valleys of 5.6 Ma (House and others, 2008). Farther south in the Blythe basin, the upper part of the Bouse Formation is 4.8 Ma based on an interbedded ash (Sarna-Wojcicki and others, 2011; Harvey, 2014). The base of the formation in the Blythe basin may be as old as 6 Ma based on the last occurrences of marine foraminifera species (McDougall and Martinez, 2014). The approximately 4.5 to 3.5 Ma Bullhead Alluvium records the deposition of fluvial gravel, sand, and minor mud during aggradation of a >200-meter-thick aggradation package that followed the establishment of a graded river profile between the Grand Wash trough and the proto-Gulf of California (House and others, 2005; Howard and others, 2015). The Palo Verde alluvium is lithologically similar to the Bullhead Alluvium, but contains rounded petrified wood that is interpreted to be reworked out of the Bullhead Alluvium (House, 2016); the geomorphic expression of these deposits is consistent with post-Bullhead Alluvium terrace remnants. Although the age of the Palo Verde alluvium is indeterminate, its landscape position indicates an age between that of the Bullhead Alluvium and the Chemehuevi Formation. The ~70 ka Chemehuevi Formation is an approximately 100-meterthick aggradation package composed of mud and sand. Its origin has been debated, but is probably related to climatic fluctuations (Malmon and others, 2011). The Riverside alluvium is composed of relatively thin veneers over a series of cut terraces linked to brief periods of stability while the river incised the Chemehuevi fill and underlying strata (House, 2016). Finally, the Blythe alluvium forms the Holocene to modern flood plain and subaqueous sediments of the Colorado River channel bed (House, 2016).

One of our main goals is the mapping of these units and the adjacent piedmont deposits in the LOCO area. We are conducting new mapping and incorporating existing mapping into the newly defined stratigraphic framework. Our initial efforts have focused on the LOCO strata in the area between Hoover and Imperial Dams, although some mapping in the Lake Mead area and south of the Chocolate Mountains was also either conducted or compiled. The digital database presented here contains new mapping and compilation efforts to date (fig. 1). Future work will focus on the incorporation of the Riverside alluvium. We are also in the process of using this database to test hypotheses about post-6 Ma deformation in the area and the depositional processes of the Bouse Formation. Here we present only the new mapping and mapping compilation for other researchers and stakeholders to use as our study continues.

Methods

This section summarizes the methods used in the map compilation, which are explained in more detail in the metadata. Geologic units were extracted from published and unpublished mapping, including map unit polygons from in-progress USGS and AZGS mapping campaigns (see the metadata for a complete list of map references). We only incorporated mapping at scales of 1:150,000 or greater (more detailed) and created separate feature classes to aggregate map units by original mapping scale. The source dataset of each polygon was recorded, as well as the original map unit designation. We then assigned each polygon to a stratigraphic unit—Bouse Formation, Bullhead Alluvium, Palo Verde alluvium, Chemehuevi Formation, or Blythe alluvium—in our new stratigraphic framework (House, 2016). When possible, these units were further subdivided based on facies. In some cases, previous mapping was reassigned to other stratigraphic units based on field observations or analytical data. For example, certain outcrops of the Muddy Creek deposits in the western Lake Mead area were here assigned to the Bouse Formation(?) based on the presence of the 5.6 Ma ash of Wolverine Creek (Castor and Faulds, 2001) and ⁸⁷Sr/⁸⁶Sr data (Roskowski and others, 2010). Where pre-existing mapping was known to include multiple stratigraphic units of interest, those polygons were assigned to the dominant unit and the map unit was queried. We also noted when the units of interest were interbedded or grouped with piedmont deposits. During compilation, where mapping overlapped, we excluded (1) older mapping that did not benefit from recent advances in the understanding of Colorado River stratigraphy, (2) mapping that was less detailed, (3) incomplete mapping (even if it was more detailed), and (4) duplicate polygons of the same geometry. Polygons sometimes had to be split at map boundaries so that only part of a polygon could be included in the compilation. Our resulting compilation contains a single polygon feature class for each map unit, derived from the most accurate and detailed mapping. Consequently, the original scale of map units in our compilation will vary spatially. Again, we plan to update this product at regular intervals in conjunction with ongoing mapping.

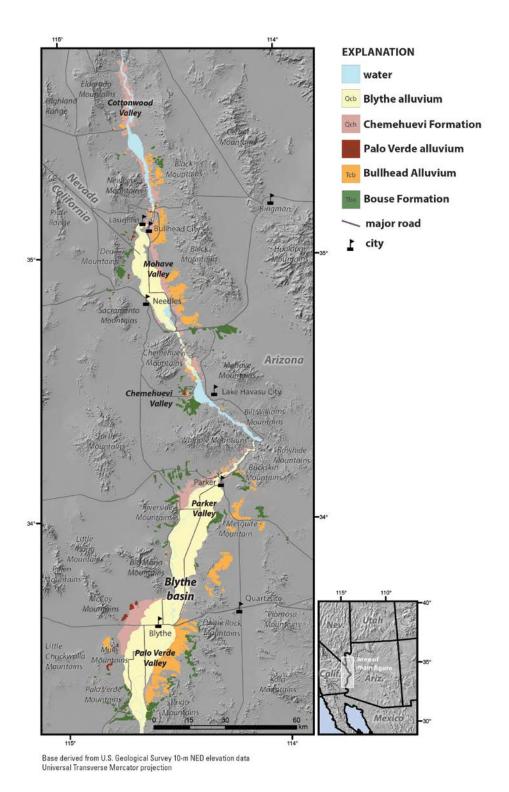


Figure 1. Map showing the study area and the distribution of select strata of the lower Colorado River corridor (LOCO) group from new and previously published mapping. Future efforts will focus on the incorporation of the Riverside alluvium (post-Chemehuevi Formation), continued new mapping, and remapping where needed.

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