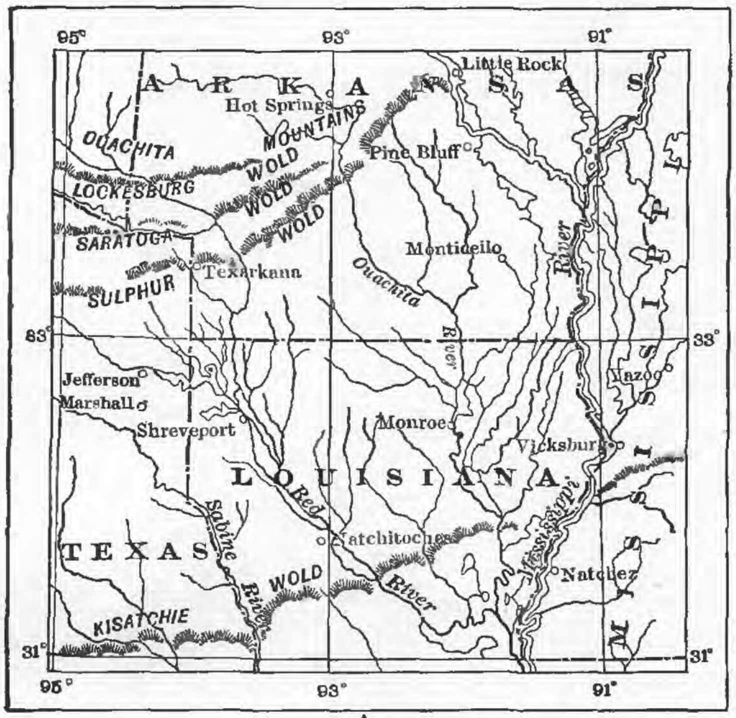
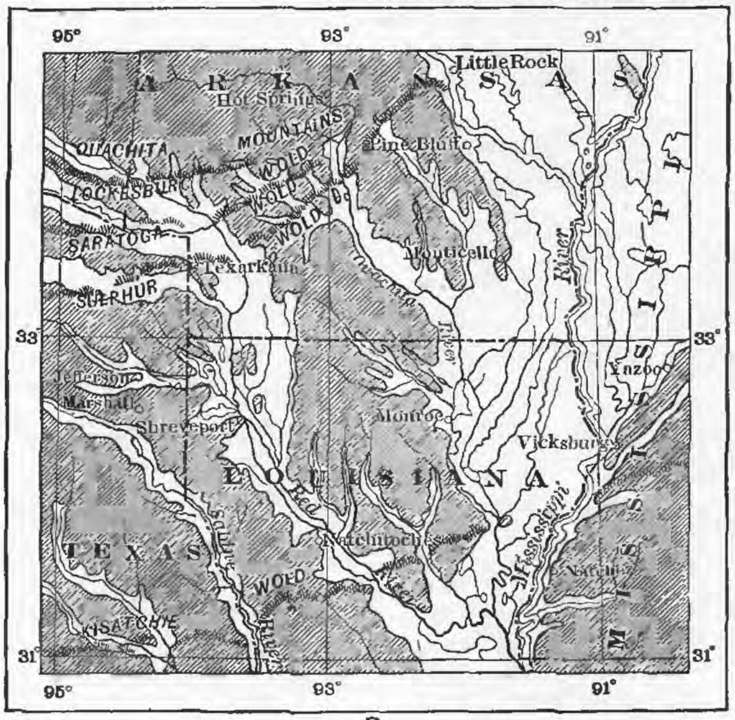


A

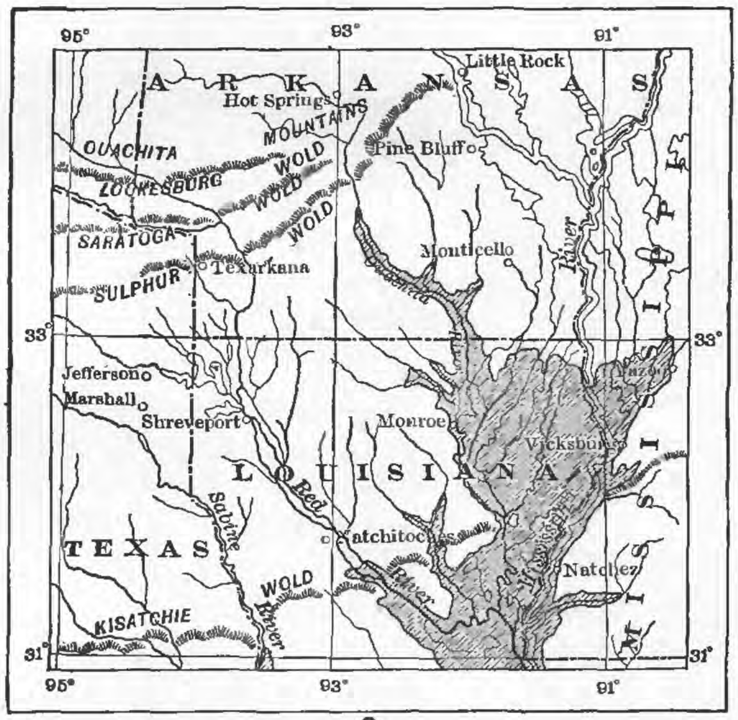
The Lockesburg, Saratoga, Sulphur, and Kisatchie wolds.



B

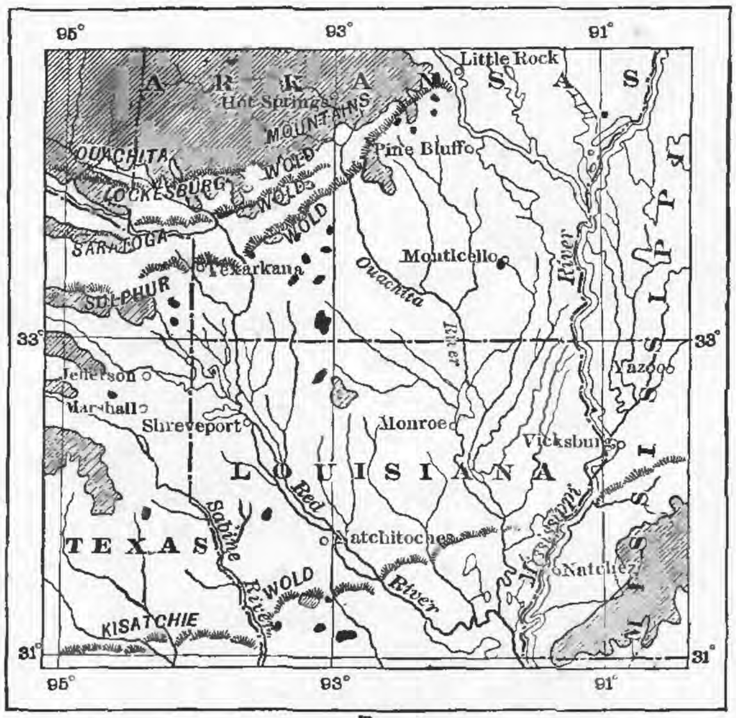
Hill lands.

Flood-plain and terrace areas.



C

Areas less than 100 feet above gulf level.



D

Areas more than 400 feet above gulf level.

SALIENT TOPOGRAPHIC FEATURES OF THE GULF COASTAL PLAIN IN NORTHERN LOUISIANA AND SOUTHERN ARKANSAS.

BY

A.C.VEATCH.

Scale

0 50 100 200 300 miles

1905

Arkansas-Louisiana State line (fig. 1). These lakes are the most important recent topographic features of this region, having been formed since the fifteenth century, but now the cause of their formation having passed, they are returning to their normal status as tributary streams. Several no longer exist, though still represented on maps because of the lack of recent detailed surveys (p. 61).

In the hill lands the general character of the topography is irregular and rolling, the hills rising 100 to 200 feet above the flat-bottomed stream channels which extend in every direction, but the unequal hardness of the underlying beds has given rise

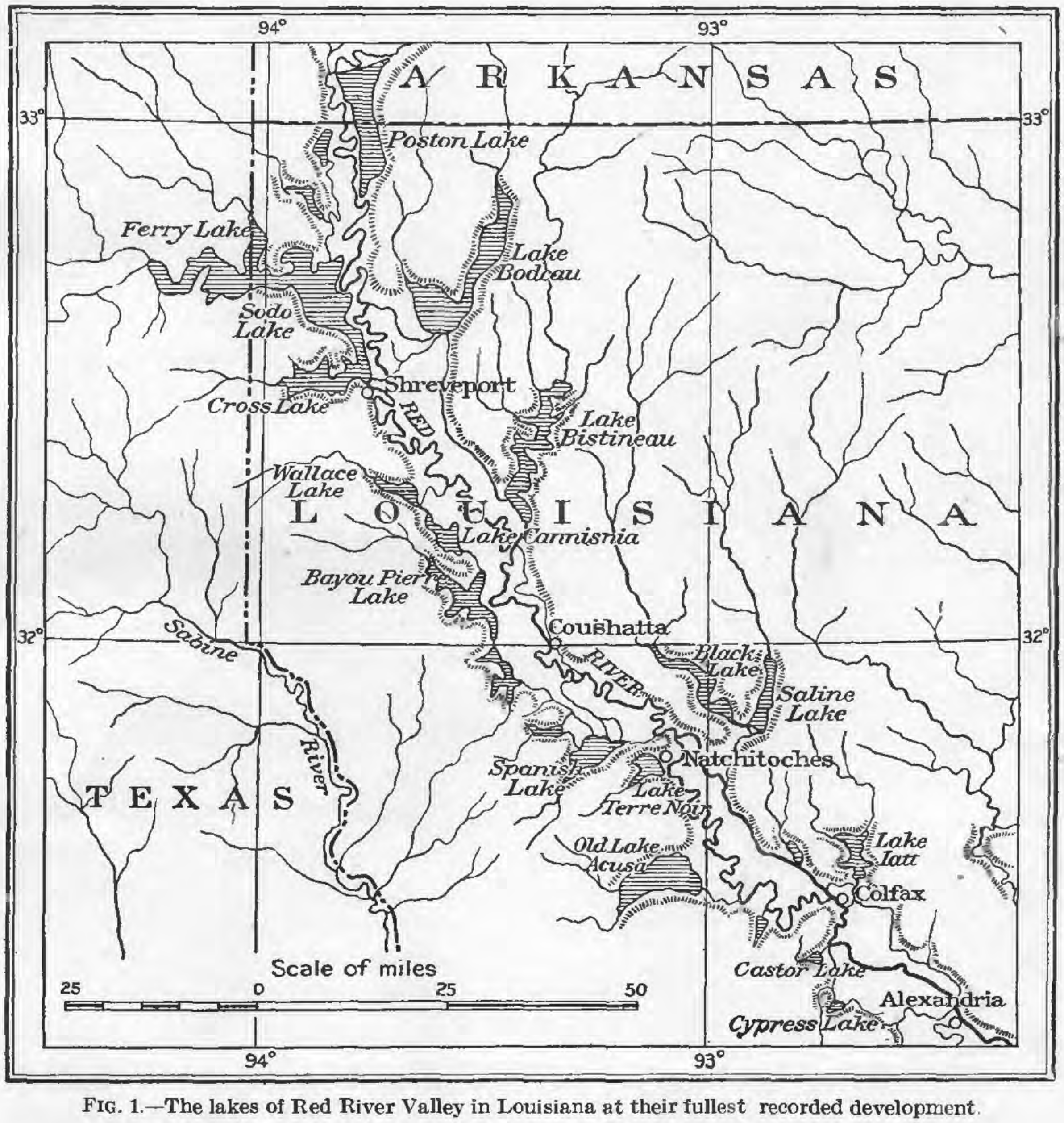


FIG. 1.—The lakes of Red River Valley in Louisiana at their fullest recorded development.

to several transverse ranges of hills, which are more or less persistent for many miles and follow the general strike of the formations producing them. Of these the Kisatchie Wold[a] (Pl. I, *A*), which is produced by the hard sandstone layers in the Catahoula (Grand Gulf) formation, is perhaps the most important. Others are the Sulphur Wold, formed by the sandy beds of the lower Eocene, and the Saratoga and Locksburg wolds, by Cretaceous formations. The transverse valley or vale[a] to

[a] For definition and derivation of the terms wold and vale see Prof. Paper U. S. Geol. Survey No. 44, 1906, p. 29.

CRETACEOUS.

CONDITIONS OF DEPOSITION.

The depression just mentioned marked the beginning of the Cretaceous, and in the ocean which then covered this part of North America (Pl. II, *A*) all the deposits which are now recognized as the main formations of this portion of the Coastal Plain were laid down in Cretaceous and Tertiary time. The alternations of sands, clays, marls, and limestones which now underlie this region indicate varying depths of ocean water on this old sea floor, and show that many changes took place in the relative elevation of the land and water, but the submergence beneath the sea was essentially continuous from early Cretaceous to late Oligocene time.

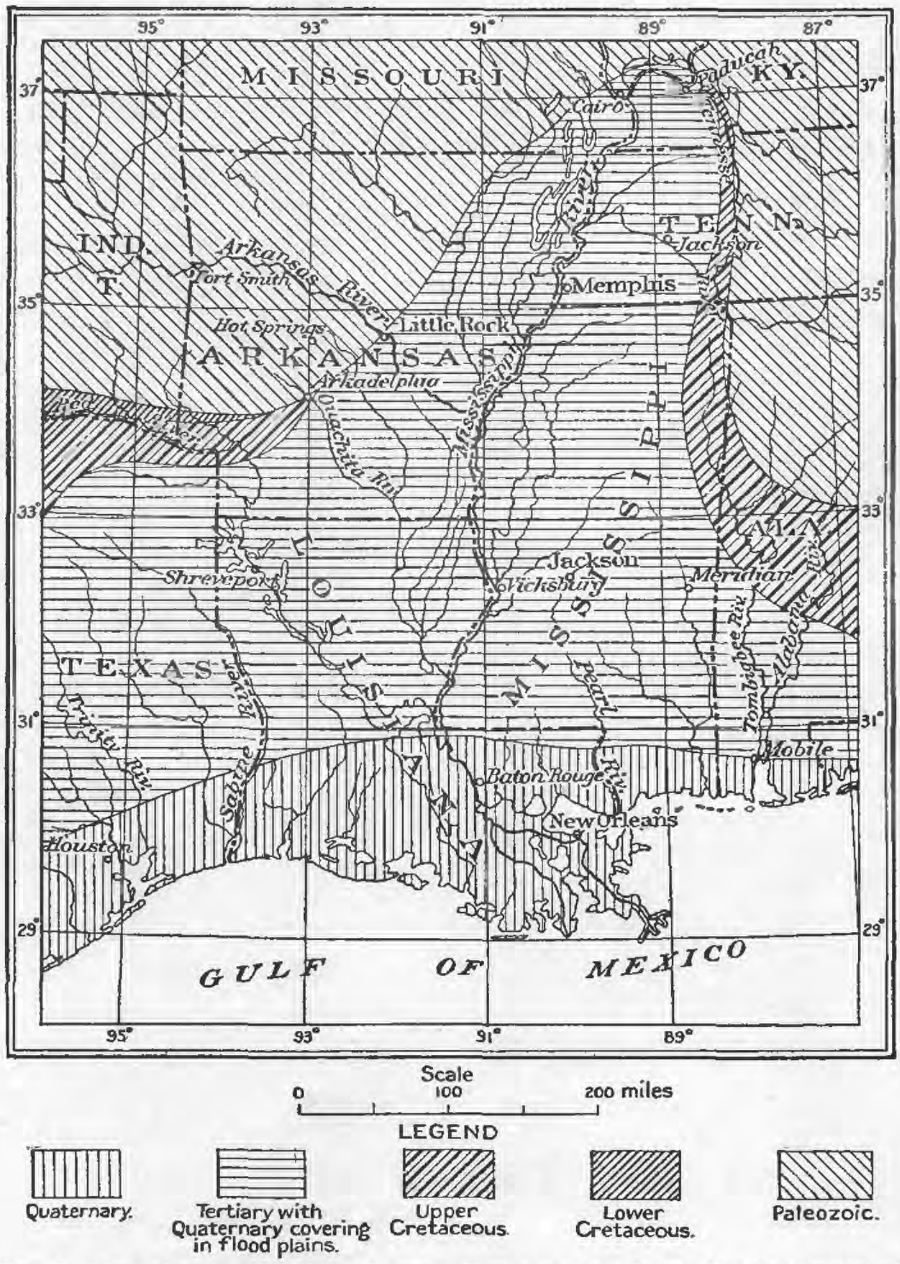
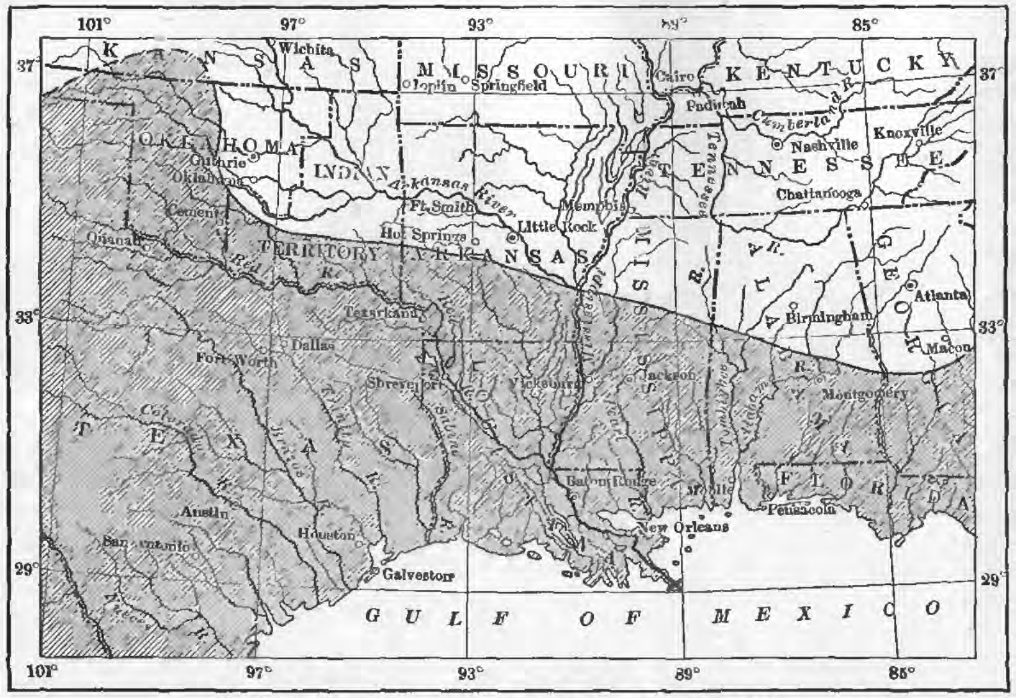


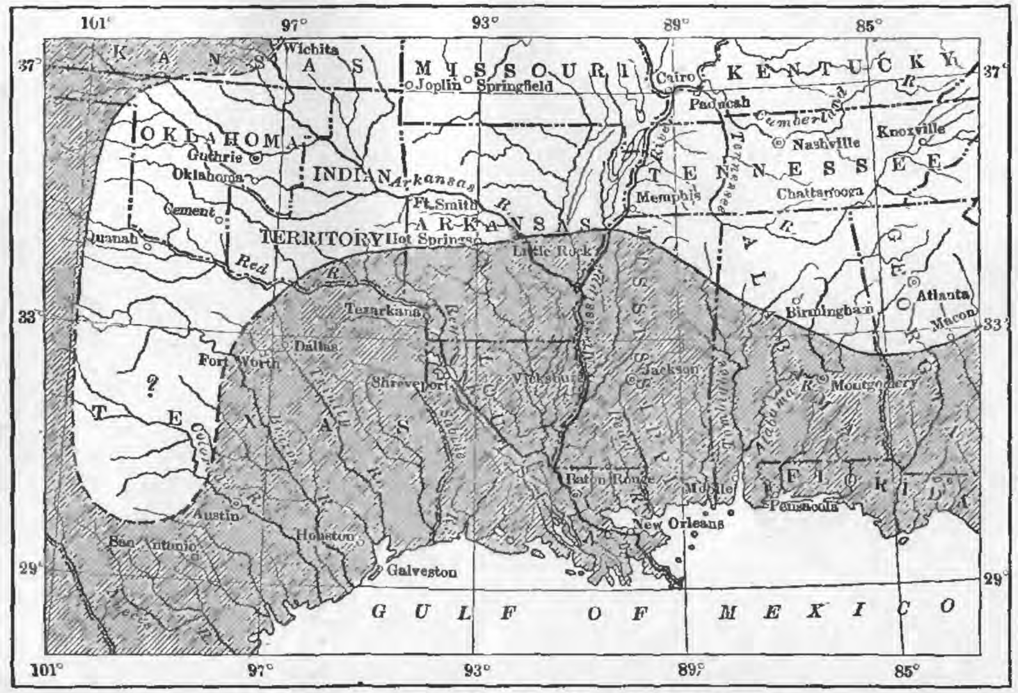
FIG. 3.—Map showing overlap of upper Cretaceous on lower Cretaceous and of Tertiary on Cretaceous in the Mississippi Valley.

The different formations recognized in this region represent periods of deposition during which the conditions were essentially the same and certain forms of life reached a peculiarly characteristic development. The progressive changes in the character of the animals inhabiting these waters make their remains—the fossil shells and bones so common over much of the region—of great value as indicators of the time at which the particular beds in which they occur were deposited and, therefore, the position of these beds in the general series of deposits.

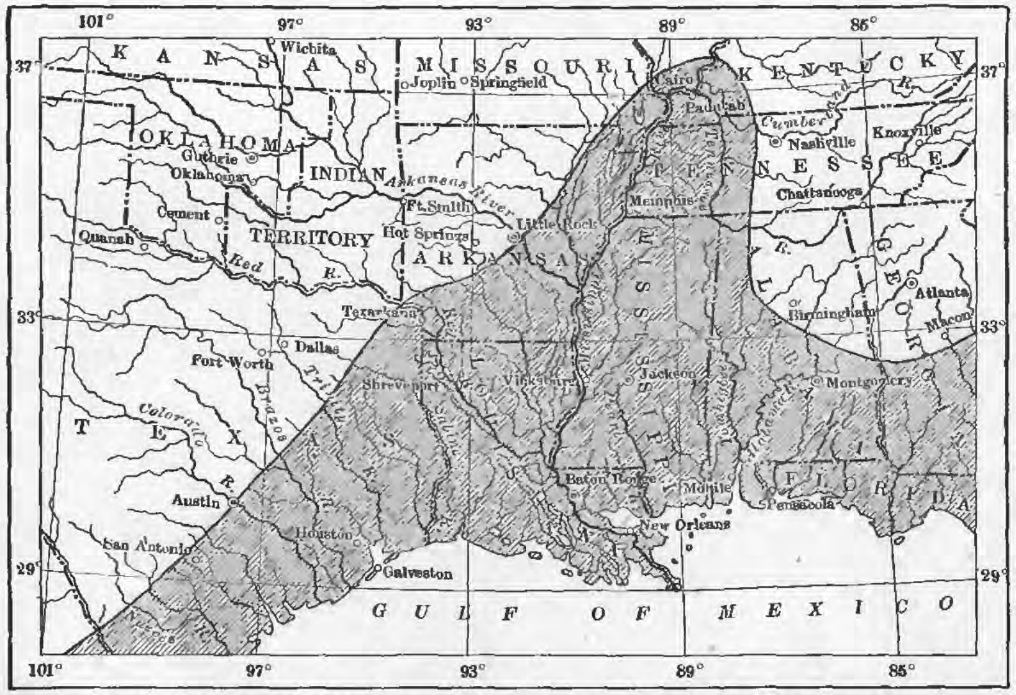
If these fluctuations in the relative elevation of the land and sea had occurred entirely in one plane, that is, if the land had been lifted and lowered without the shore line of one period cutting that of another, the present outcrops of the deposits would have been roughly parallel. Since, however, these movements were irregular and there was warping in a direction not entirely parallel to the first shore line, the



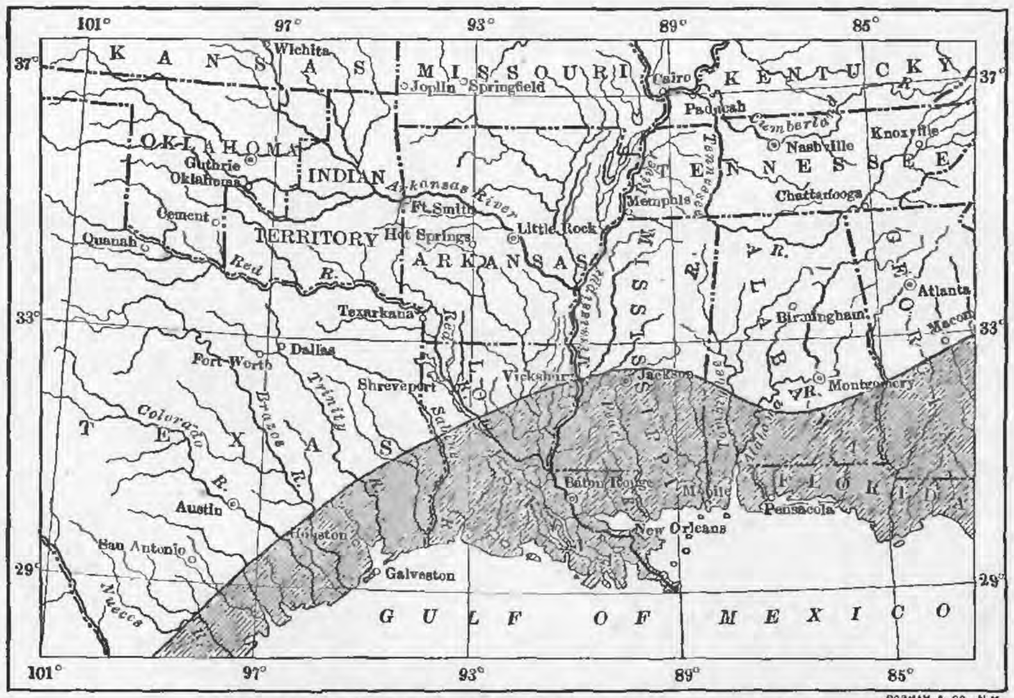
(A) EARLY LOWER CRETACEOUS



(B) EARLY UPPER CRETACEOUS



(C) LATE CRETACEOUS AND EARLY TERTIARY



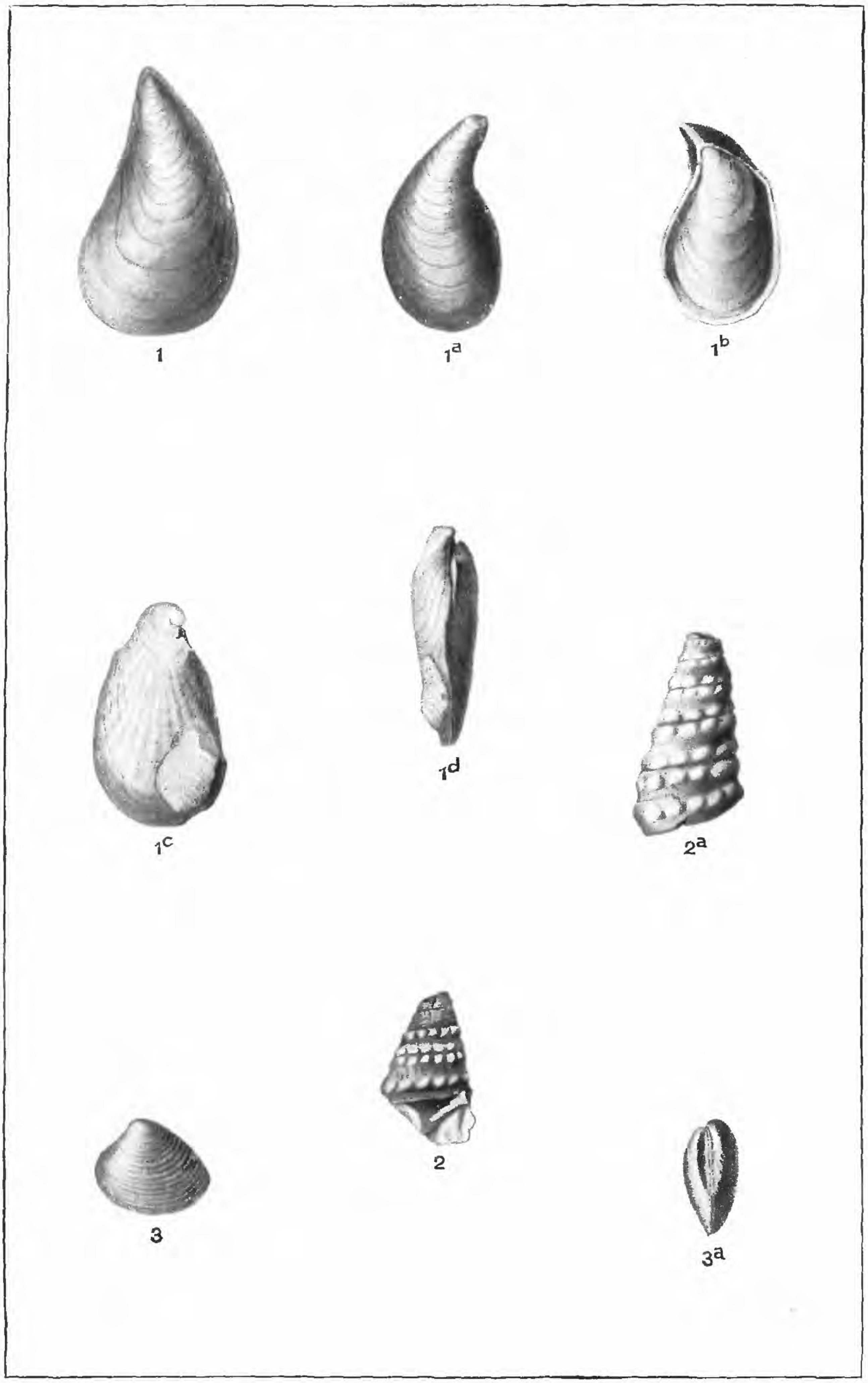
(D) EARLY OLIGOCENE TERTIARY

BORMAY & CO., N.Y.

Scale
100 0 100 200 300 400 500 600 MILES

RELATIVE POSITIONS OF LAND AND WATER AREAS IN THE SOUTH CENTRAL UNITED STATES DURING THE CRETACEOUS AND TERTIARY EPOCHS.

Water, present Gulf of Mexico and all shaded areas; land, unshaded portion of present land.



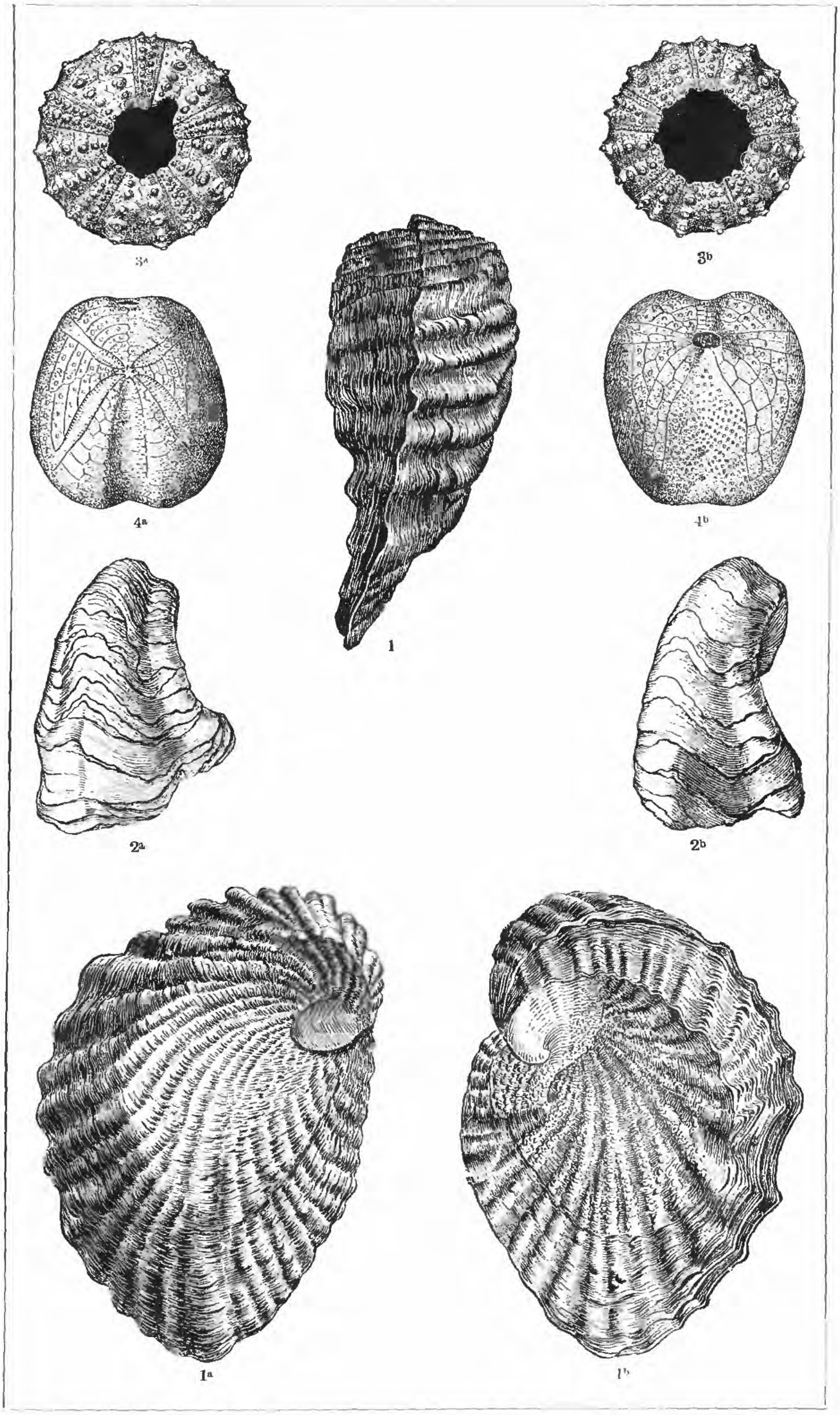
CHARACTERISTIC FOSSILS OF THE TRINITY FORMATION IN SOUTHERN ARKANSAS.

1, 1a, 1b, 1c, 1d. Ostrea franklini Coquand. 2, 2a. Glauconia branneri Hill. 3, 3a. Astarte? pikensis Hill.

A. GYPSUM BED AT PLASTER BLUFF, ON LITTLE MISSOURI RIVER, 3 MILES SOUTH OF MURFREESBORO, ARK.

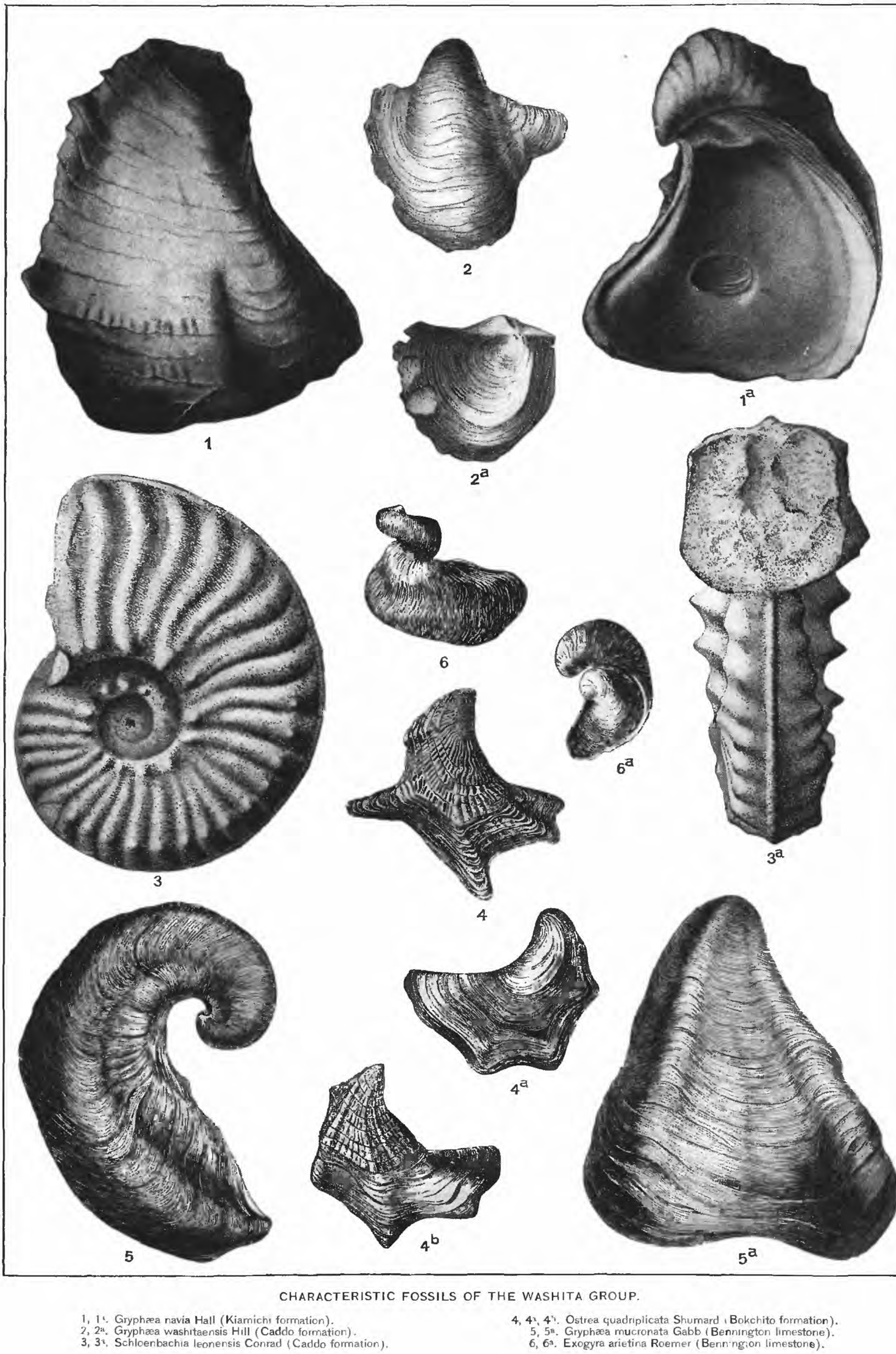
An outcrop of the Trinity formation. Photograph by J. A. Taff.

B. TYPICAL LANDSLIP BENCH, SIDE OF ROCKY MOUNTAIN, 5 MILES EAST OF HOPE, ARK.



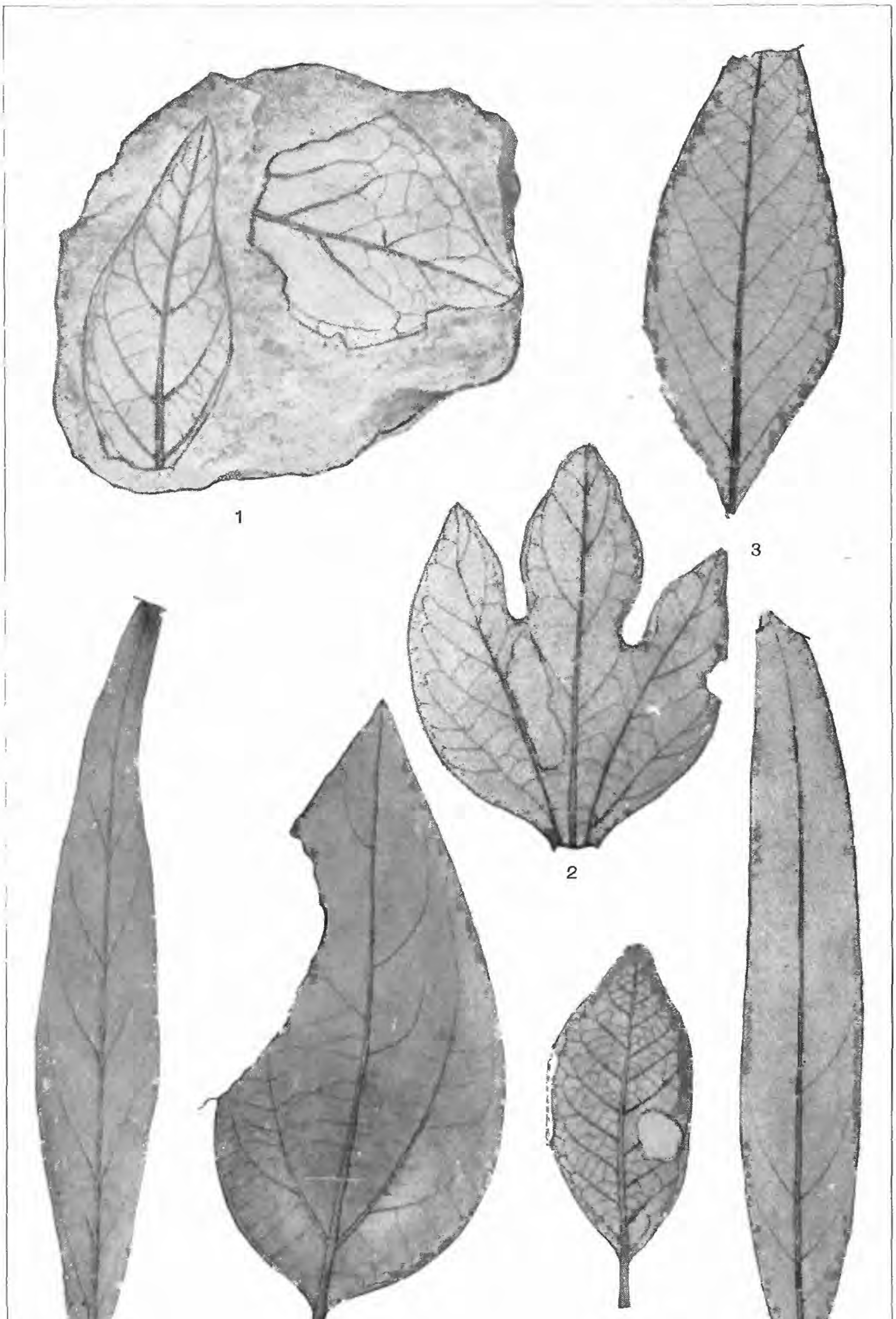
CHARACTERISTIC FOSSILS OF THE GOODLAND LIMESTONE.

1, 1a, 1b. Exogyra texana Roemer.
2a, 2b. Gryphæa marcoui Hill and Vaughan.
3a, 3b. Pseudodiadema texanum Roemer.
4a, 4b. Enallaster texanus Roemer.



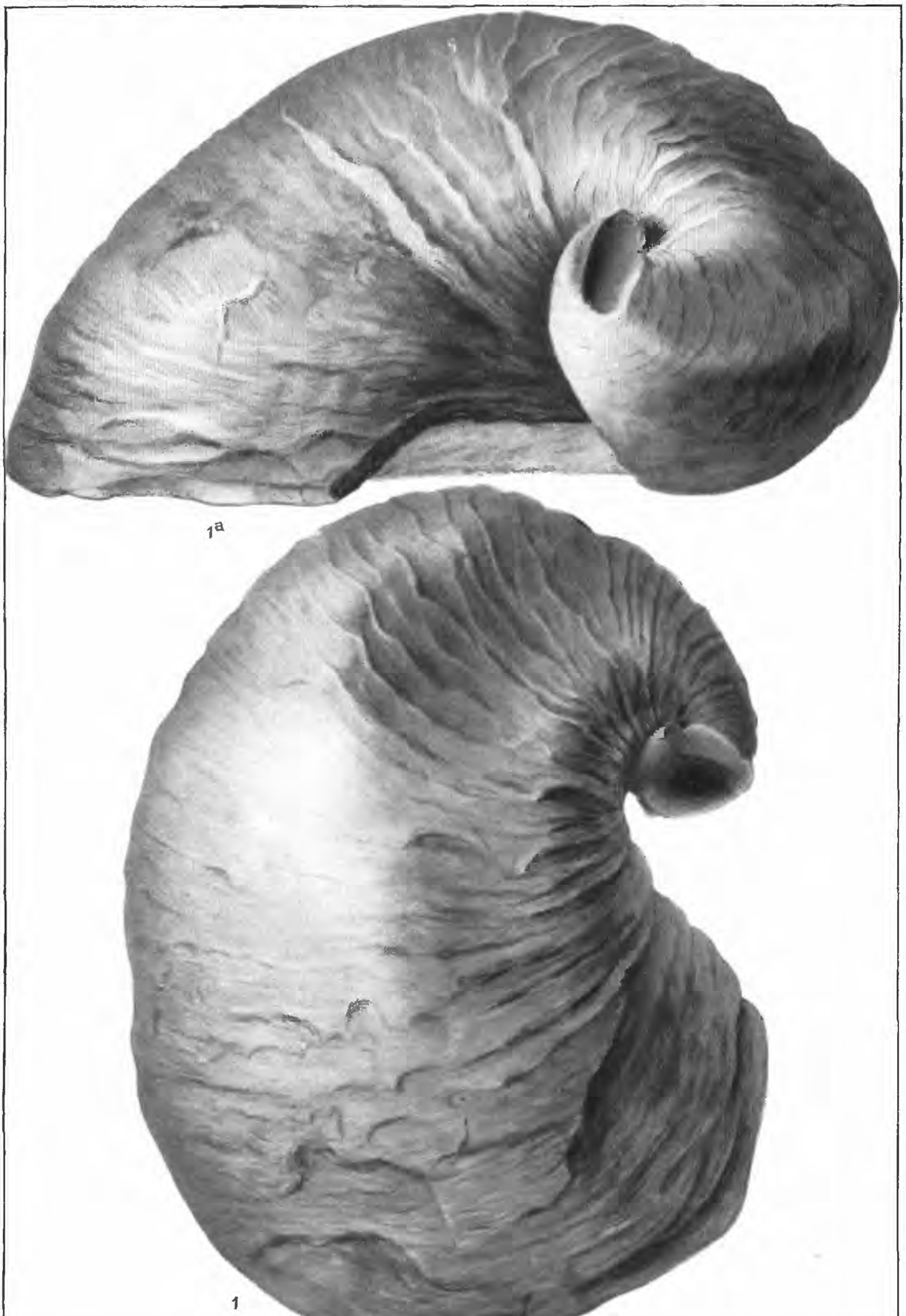
CHARACTERISTIC FOSSILS OF THE WASHITA GROUP.

1, 1a. Gryphæa navia Hall (Kiamichi formation).
2, 2a. Gryphæa washitaensis Hill (Caddo formation).
3, 3a. Schloenbachia leonensis Conrad (Caddo formation).
4, 4a, 4b. Ostrea quadriplicata Shumard (Bokchito formation).
5, 5a. Gryphæa mucronata Gabb (Bennington limestone).
6, 6a. Exogyra arietina Roemer (Bennington limestone).



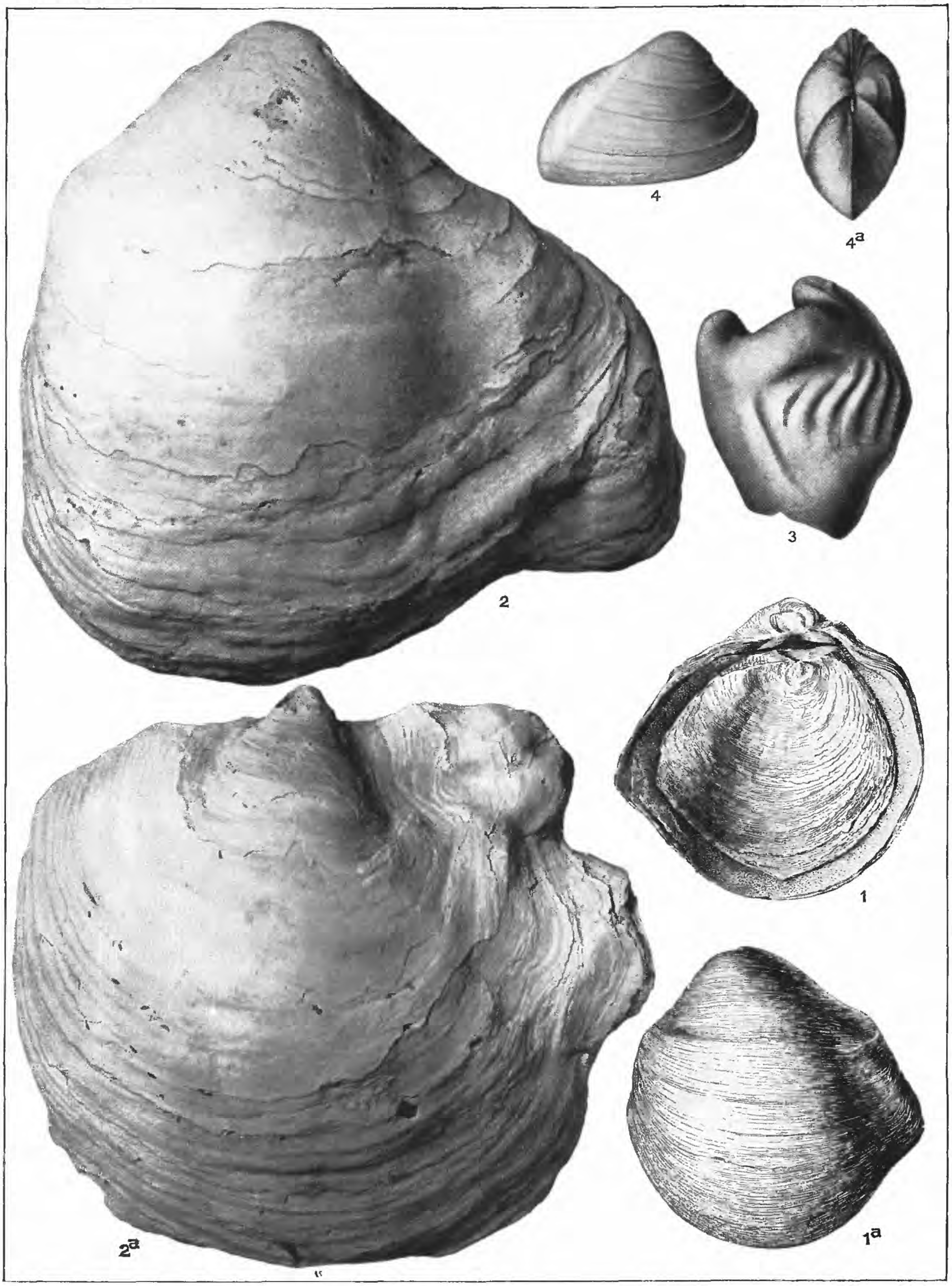
FOSSIL LEAVES FROM THE WOODBINE FORMATION.

1. Diospyros steenstrupi? Heer.
2. Lindera venusta Lesq.
3. Diospyros primæva Heer.
4. Andromeda pfaffiana Heer.
5. Cinnamomum heeri Lesq.
6. Salix hayei Lesq.
7. Myrica longa Heer.



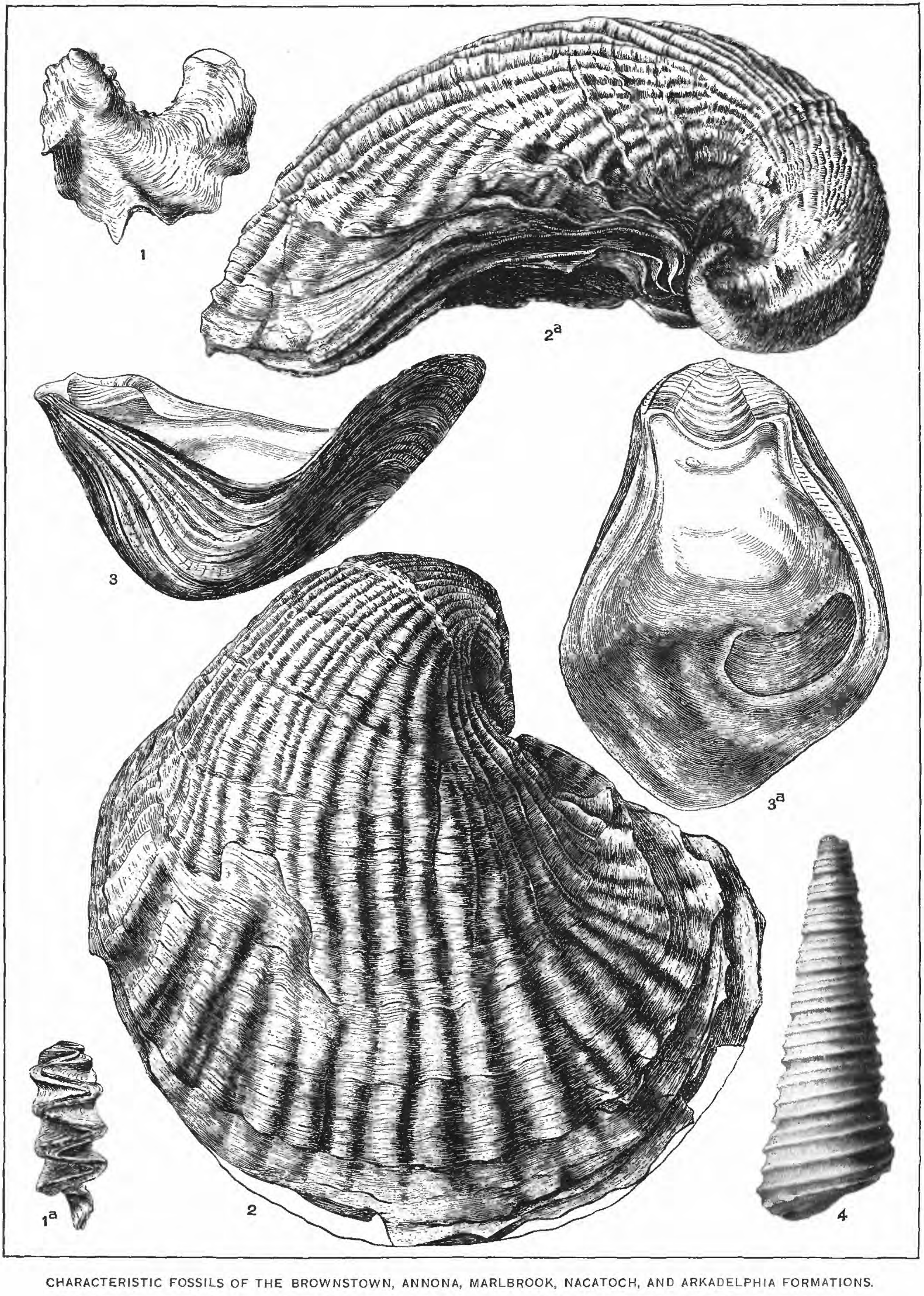
THE CHARACTERISTIC FOSSIL OF THE BROWNSTOWN FORMATION.

1, 1a. *Exogyra ponderosa* Roemer (× 5/6).



CHARACTERISTIC FOSSILS OF THE BROWNSTOWN, ANNONA, MARLBROOK, NACATOCH, AND ARKADELPHIA FORMATIONS.

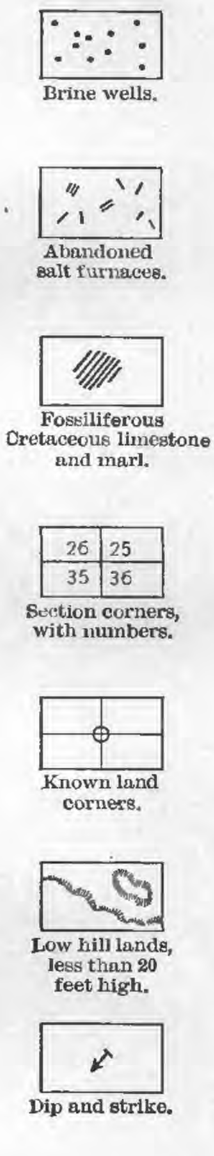
1, 1a. Gryphæa vesicularis Lamarck. This small form is common in the upper part of the Brownstown formation, and the species is found in all the formations mentioned.
2, 2a. Gryphæa vesicularis. This large form is characteristic, in Arkansas, of the base of the Saratoga chalk or mid-Marlbrook formation.
3. Pugnellus densatus Conrad.
4, 4a. Veniella lineata (Shumard).

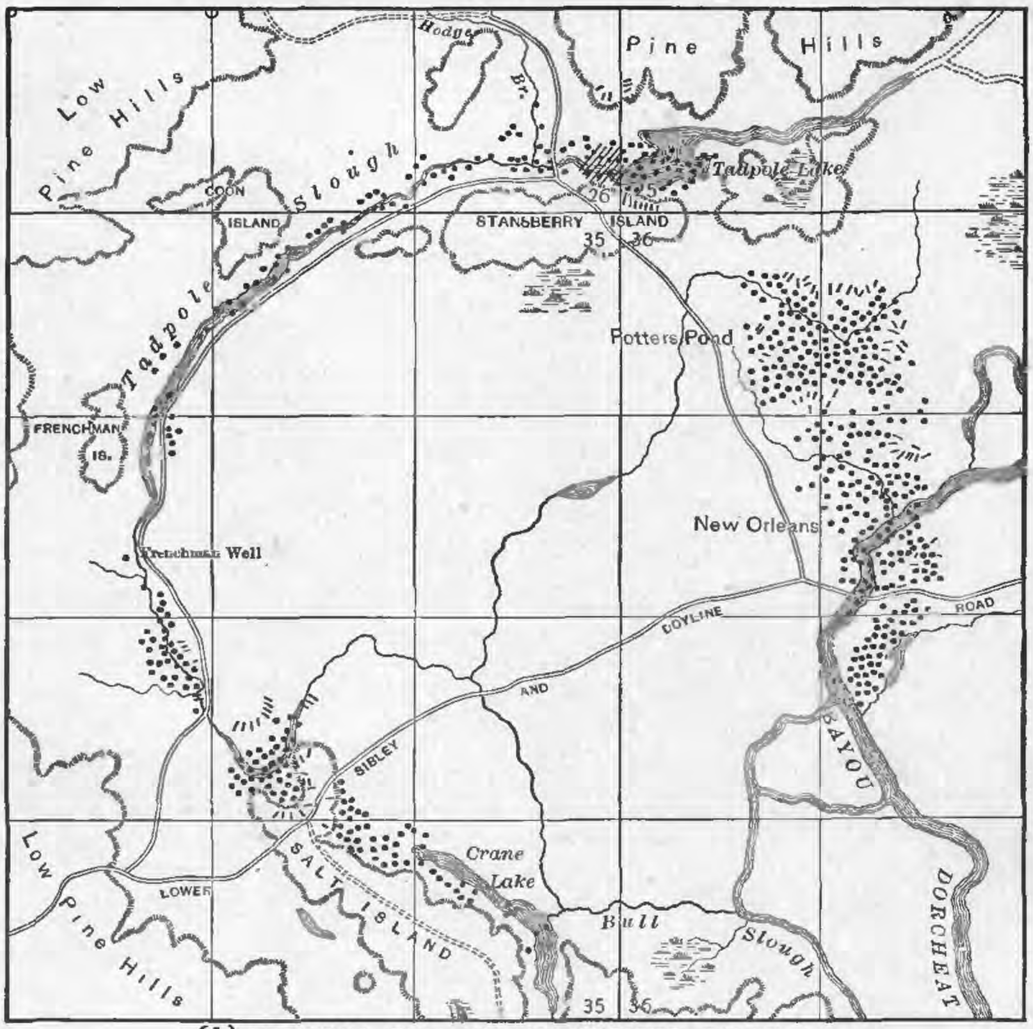


CHARACTERISTIC FOSSILS OF THE BROWNSTOWN, ANNONA, MARLBROOK, NACATOCH, AND ARKADELPHIA FORMATIONS.

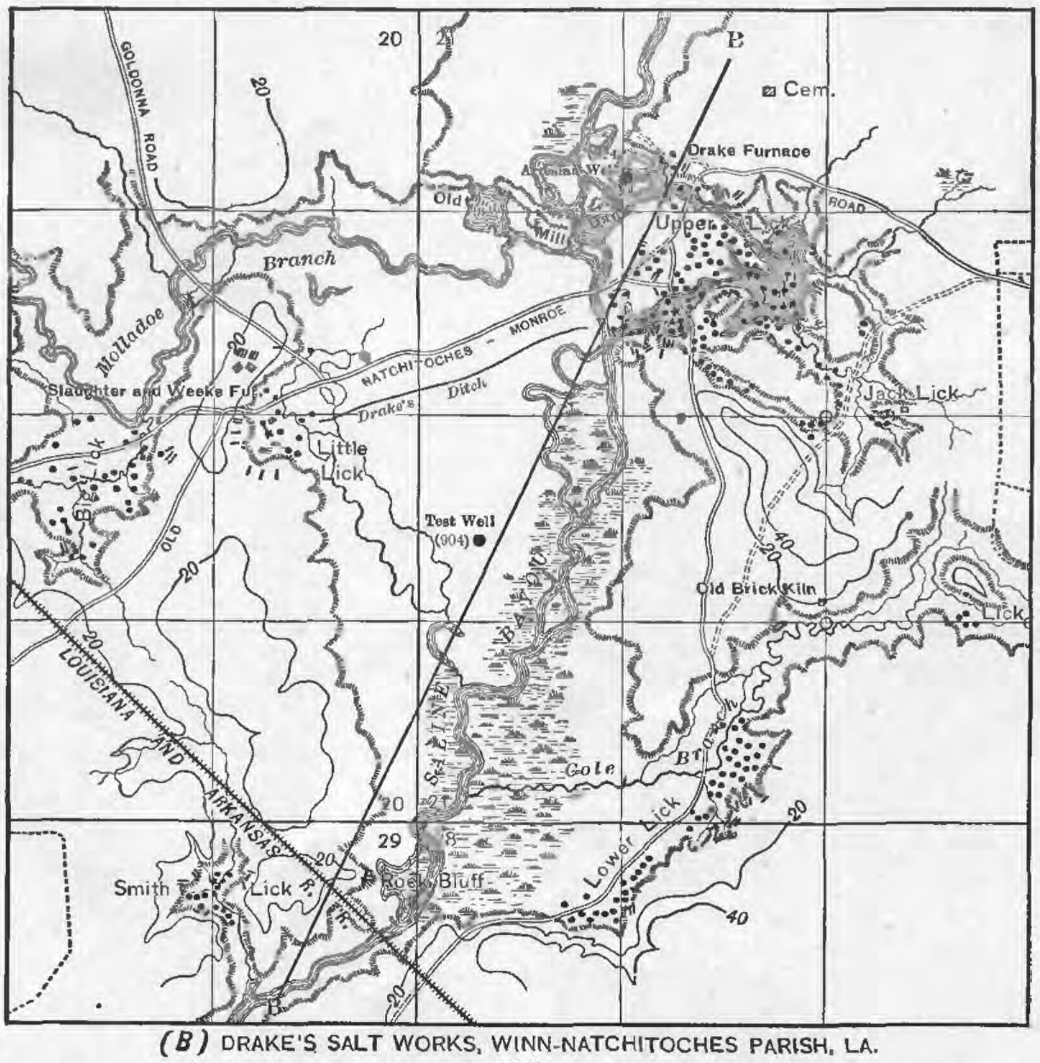
1, 1a. Ostrea larva Lamarck (Brownstown, Marlbrook, Nacatoch, Arkadelphia formations).
2, 2a. Exogyra costata Say (Marlbrook, Nacatoch, and Arkadelphia formations).
3, 3a. Ostrea subspatulata Forbes (Nacatoch and Arkadelphia formations).
4. Turritella trilira Conrad.

LEGEND





(A) BISTINEAU SALT WORKS, WEBSTER PARISH, LA.
1900.



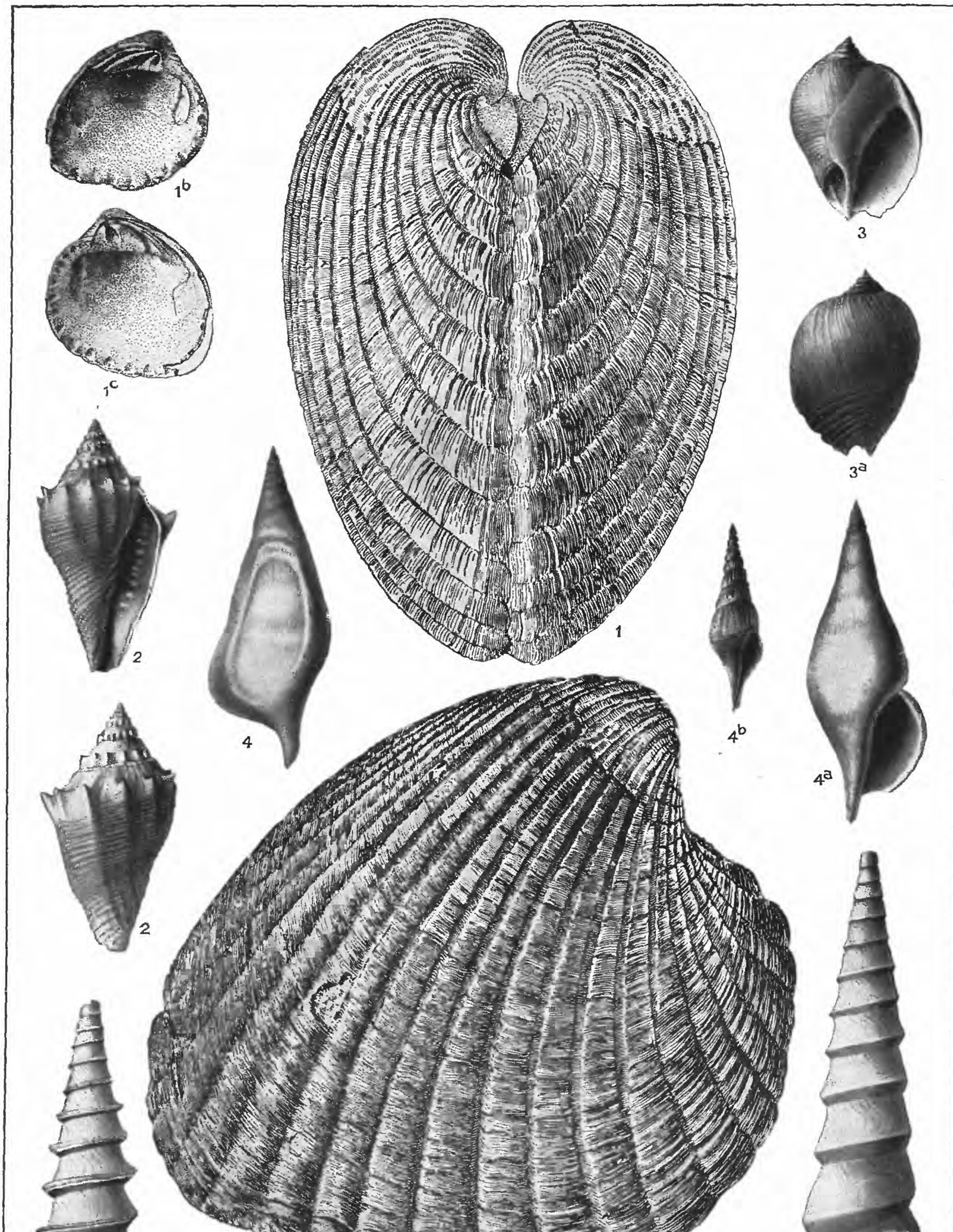
(B) DRAKE'S SALT WORKS, WINN-NATCHITOCHES PARISH, LA.
1899.

MAPS SHOWING CHARACTERISTIC, SYMMETRICAL, CIRCULAR ARRANGEMENT OF SHALLOW BRINE WELLS AROUND THE TRUNCATED DOMES OF NORTHERN LOUISIANA.

BY A. C. VEATCH

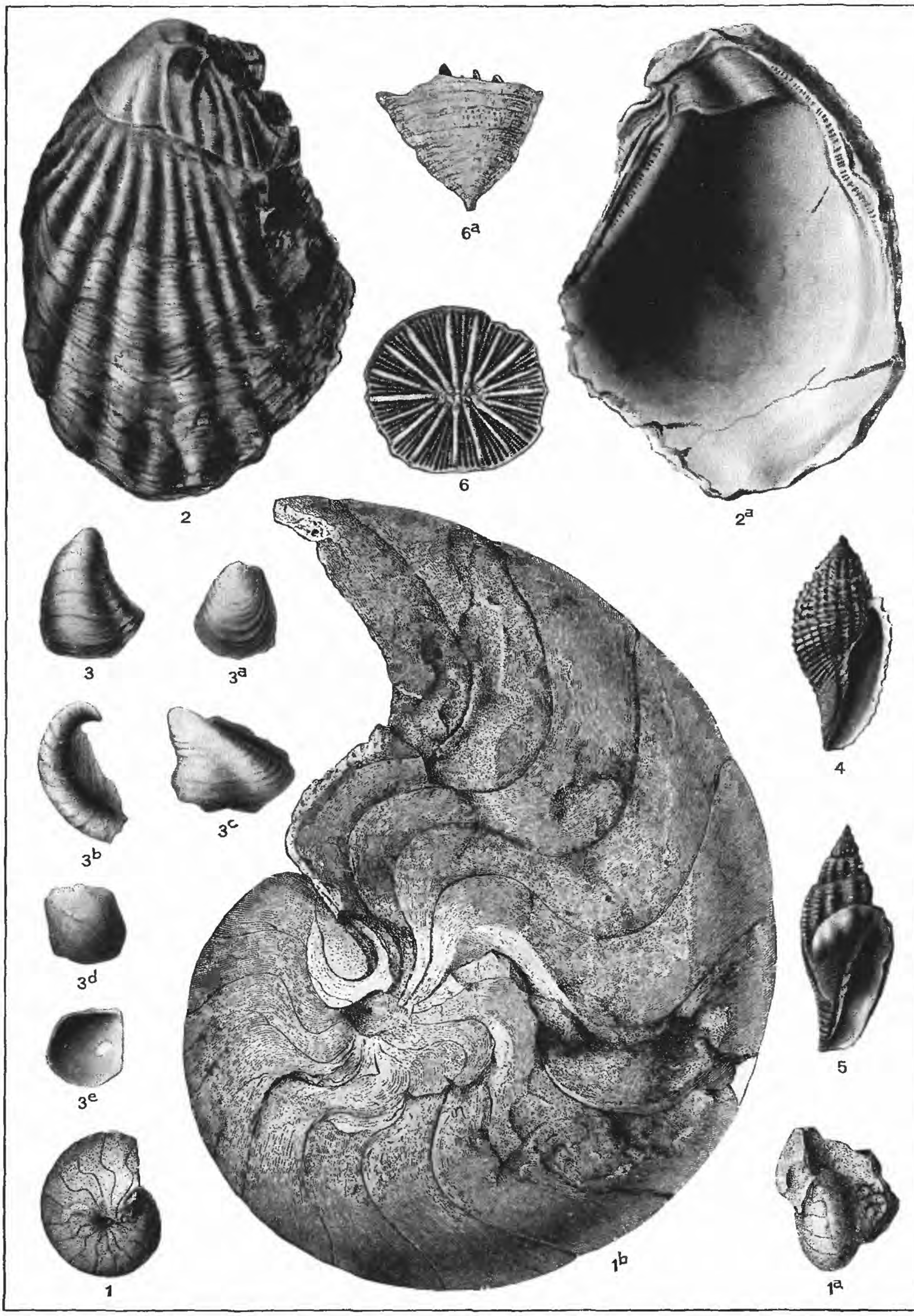
Scale

0 ¼ ½ ¾ 1 mile



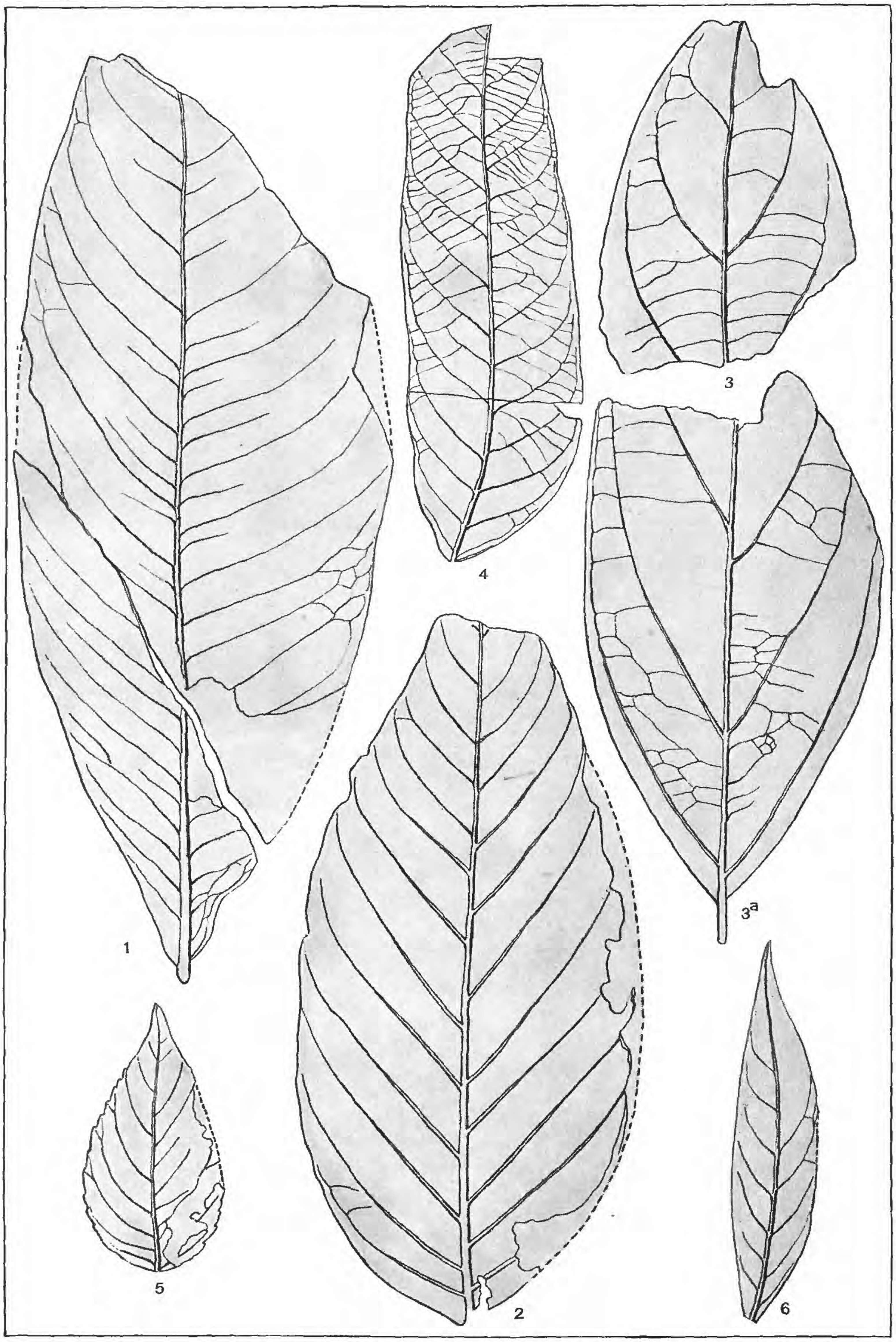
CHARACTERISTIC FOSSILS OF THE EOCENE EPOCH.

1, 1a, 1b, 1c. *Venericardia planicosta* Lamarck (Midway, Sabine, Claiborne, and Jackson formations).
2, 2a. *Volutilithes petrosus* Conrad (Sabine, Claiborne, and Jackson formations).
3, 3a. *Pseudoliva vetusta* Conrad (Midway, Sabine, Claiborne, and Jackson formations).
4, 4a, 4b. *Calyptraphorus velatus* Conrad (4b shows the *young* of this species; Midway, Sabine, Claiborne, and Jackson formations).
5, 5a. *Turritella mortoni* Conrad (Midway and Sabine formations).



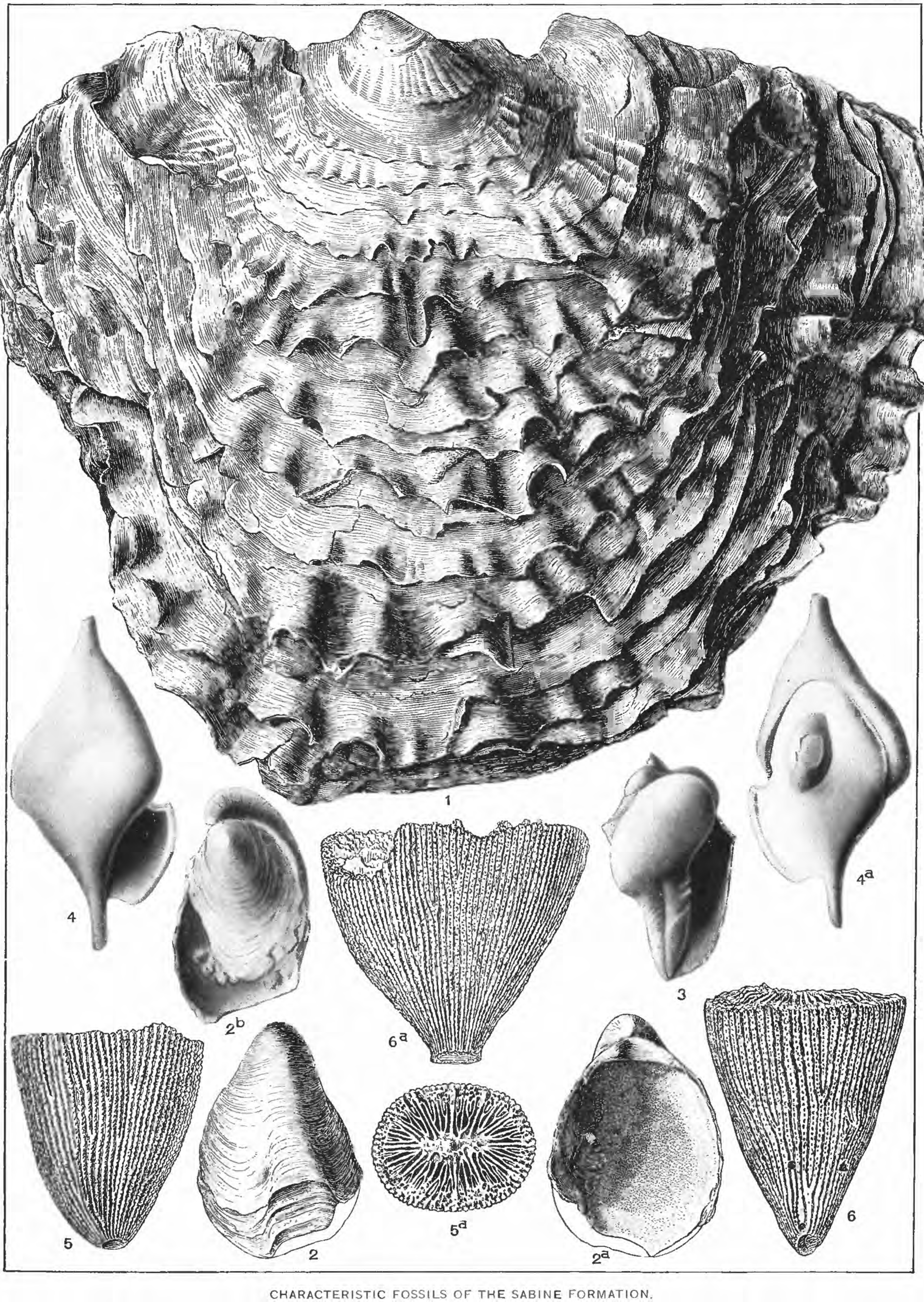
CHARACTERISTIC FOSSILS OF THE MIDWAY FORMATION.

1, 1a, 1b. Enclimatoceras ulrichi White.
2, 2a. Ostrea crenulimarginata Gabb.
3, 3a, 3b, 3c, 3d, 3e. Ostrea pulaskensis Harris.
4. Volutilithes limopsis Conrad.
5. Volutilithes rugatus Conrad.
6, 6a. Flabellum conoideum Vaughan.



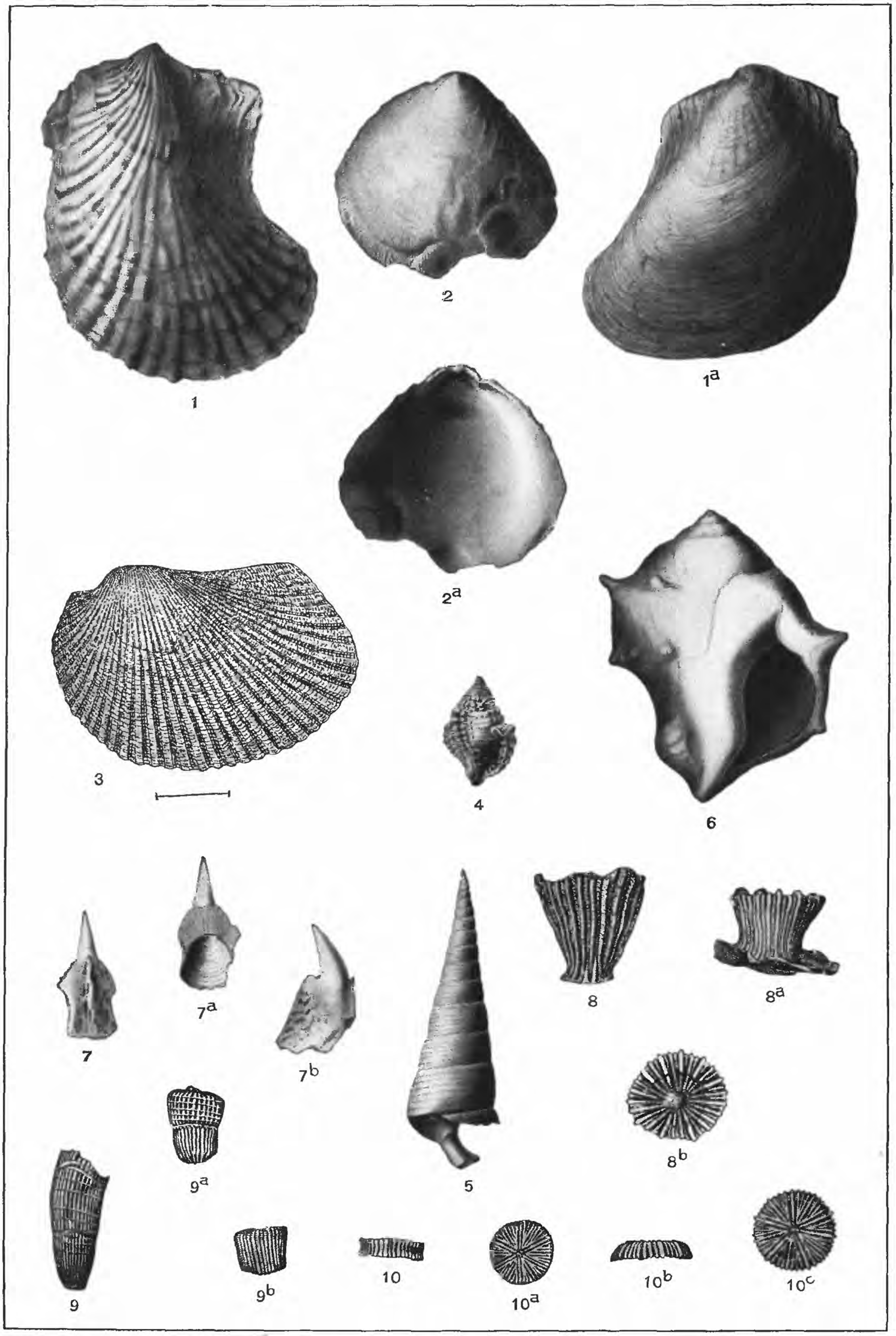
FOSSIL LEAVES FROM THE SABINE FORMATION.

1. Magnolia hilgardiana Lesquereux.
2. Andromeda eolignitica Hollick.
3, 3a. Tetranthera præcursoria Lesquereux.
4. Juglans schimperi Lesquereux.
5. Ulmus tenuinervis Lesquereux.
6. Sapindus angustifolius Lesquereux.



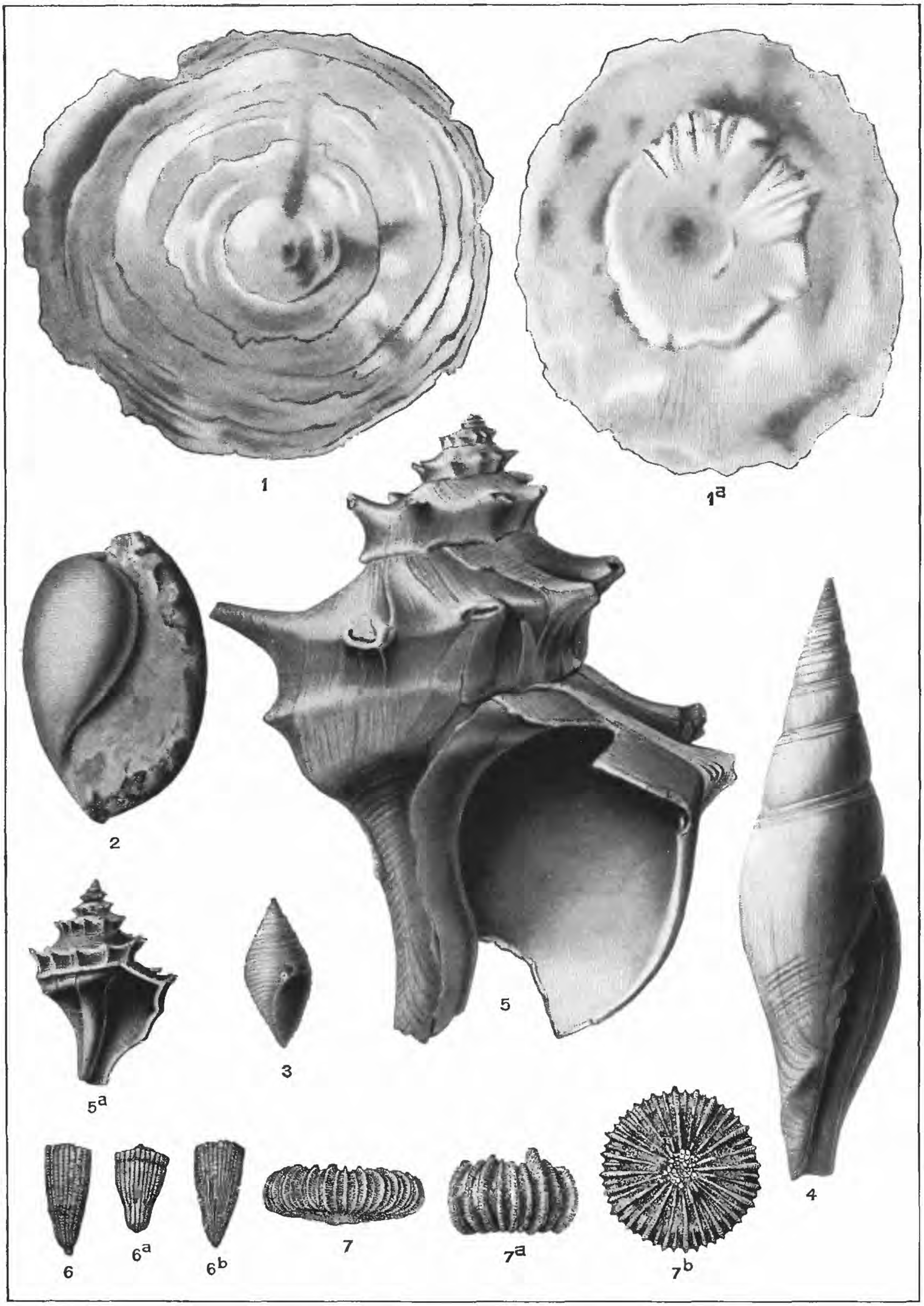
CHARACTERISTIC FOSSILS OF THE SABINE FORMATION.

1. Ostrea compresserostra Say.
2, 2a, 2b. Ostrea thirsae Gabb
3. Volutilithes petrosus var toumeyi Conrad.
4, 4a. Calyptraphorus trinodiferus Conrad.
5, 5a. Eupsammia elaborata (Conrad).
6, 6a. Balanophyllia haleana (M. Ed. and H.).



CHARACTERISTIC FOSSILS OF THE CLAIBORNE FORMATION.

1, 1a. Ostrea sellæformis Conrad.
2, 2a. Anomia ephippoides Gabb.
3. Arca vaughani Casey.
4. Pirsonella sepdentata Gabb.
5. Mesalia claibornensis Conrad.
6. Cornulina armigera Conrad.
7, 7a, 7b. Belosepia ungula Gabb.
8, 8a, 8b. Paracyathus alternatus Vaughan.
9. 9a, 9b. Paracyathus bellus Vaughan.
10, 10a, 10b, 10c. Discotrochus orbignianus M. Edw. and H.



CHARACTERISTIC FOSSILS OF THE JACKSON FORMATION.

1, 1a. Umbrella planulata Conrad.
2. Haminea grandis Aldrich.
3. Conorbis alatoides Aldrich.
4. Mitra millingtoni Conrad.
5, 5a. Levifusus branneri Harris $\frac{4}{5}$.
6, 6a, 6b. Aldrichiella elegans Vaughan.
7, 7a, 7b. Trochocyathus lunulitiformis var. montgomeriensis Vaughan.

1. Skull restored from specimens collected by Charles Schuchert in Alabama. Length 47.5 inches. The name Zeuglodon (yoke tooth) is taken from the shape of the molar teeth. Original in United States National Museum.

2. Water-worn vertebra of the type commonly found in the Jackson strata. Length 13 inches.

3. More perfect vertebra, showing processes. Length 14.75 inches.

ZEUGLODON (*BASILOSAURUS*) *CETOIDES* OWEN.

A COMMON AND CHARACTERISTIC JACKSON EOCENE WHALE-LIKE MAMMAL.

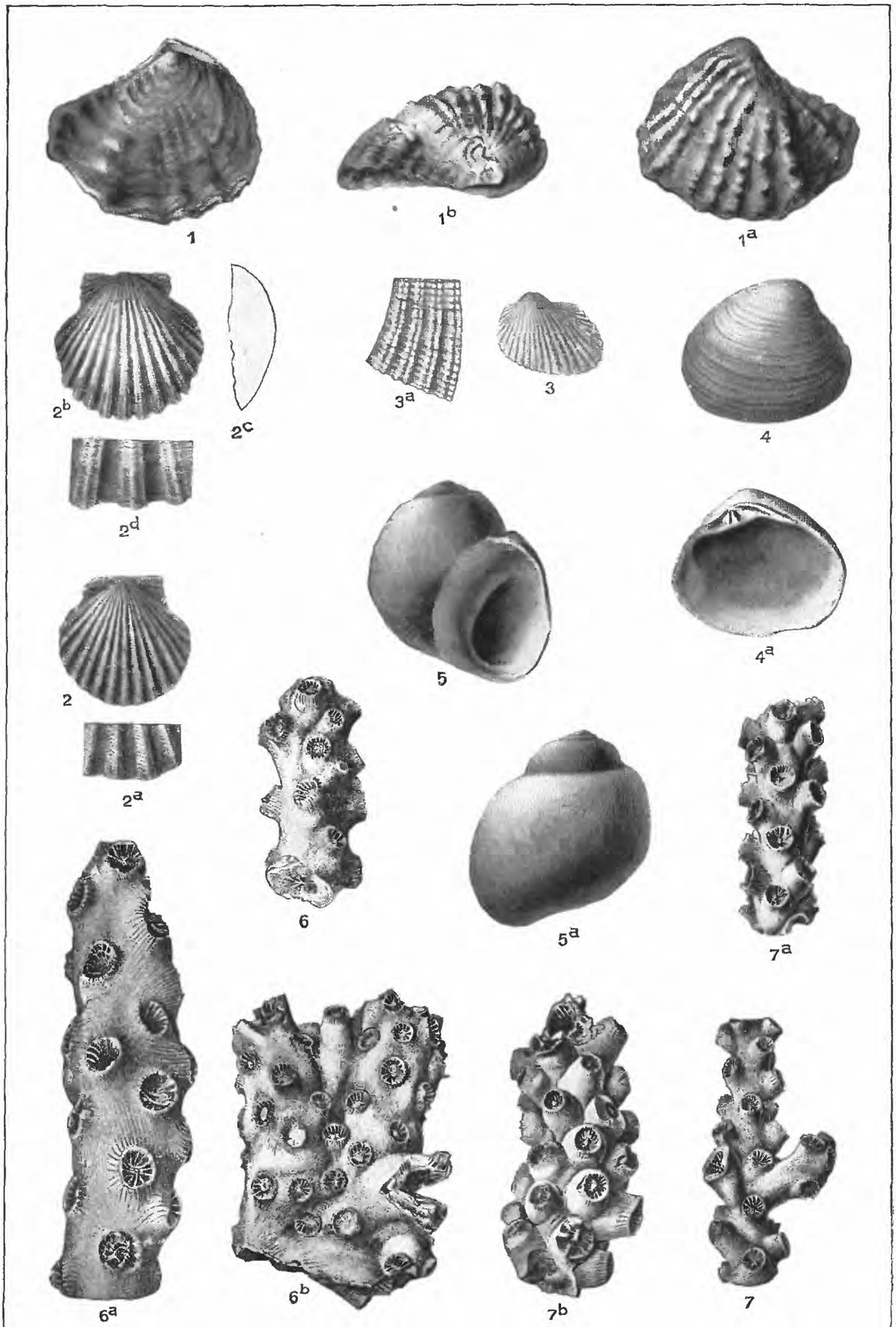
A. NOTED FOSSILIFEROUS JACKSON OUTCROP AT MONTGOMERY, LA.

Showing low southward dip of the Eocene beds. This locality is in the "steeply" dipping portion of the Caldwell-Angelina flexure and indicates the very low character of that structural feature. Photograph by G. D. Harris.

B. LOWER FALLS, MINT SPRING BAYOU, VICKSBURG, MISS.

Showing bedded limestone in the Vicksburg formation. Note unconsolidated shell marl above and below the limestone. Photograph by G. D. Harris.

NOTED OUTCROPS OF THE JACKSON AND VICKSBURG FORMATIONS.



CHARACTERISTIC FOSSILS OF THE VICKSBURG FORMATION.

1, 1a, 1b. Ostrea vicksburgensis Conrad.
2, 2a, 2b, 2c, 2d. Pecten poulsoni Morton.
3, 3a. Scapharca lesueuri Dall.
4, 4a. Macrocallista (Chionella) sobrina Conrad.
5, 5a. Ampullinopsis mississippiensis (Conrad).
6, 6a, 6b. Oculina vicksburgensis (Conrad).
7, 7a, 7b. Oculina mississippiensis (Conrad).

A. QUARRY ON CONN LEAGUE, JASPER COUNTY, TEX., NORTHEAST OF ROCKLAND.

Opened for stone for Sabine Pass jetties.

B. ON TEXAS AND PACIFIC RAILWAY, NEAR LENA, LA.

Photograph by G. D. Harris.

OUTCROPS OF THE CATAHOULA FORMATION.

A

B

LAFAYETTE GRAVEL, 5 MILES EAST OF PRESCOTT, ARK.

A. Typical exposure. *B*. Large quartzitic bowlders.

C

ESCARPMENT ON THE SOUTHERN EDGE OF THE MARKSVILLE HILLS.

A Port Hudson terrace in Avoyelles Parish, La.

OUTCROPS OF LAFAYETTE AND PORT HUDSON FORMATIONS.

In the area between the main streams the tributaries formed in time an interlocking drainage very little different from that of to-day, and at the close of this long period of erosion almost all of the old plain level had been destroyed (Pl. I, p. 14) and the major topographic features of northern Louisiana and southern

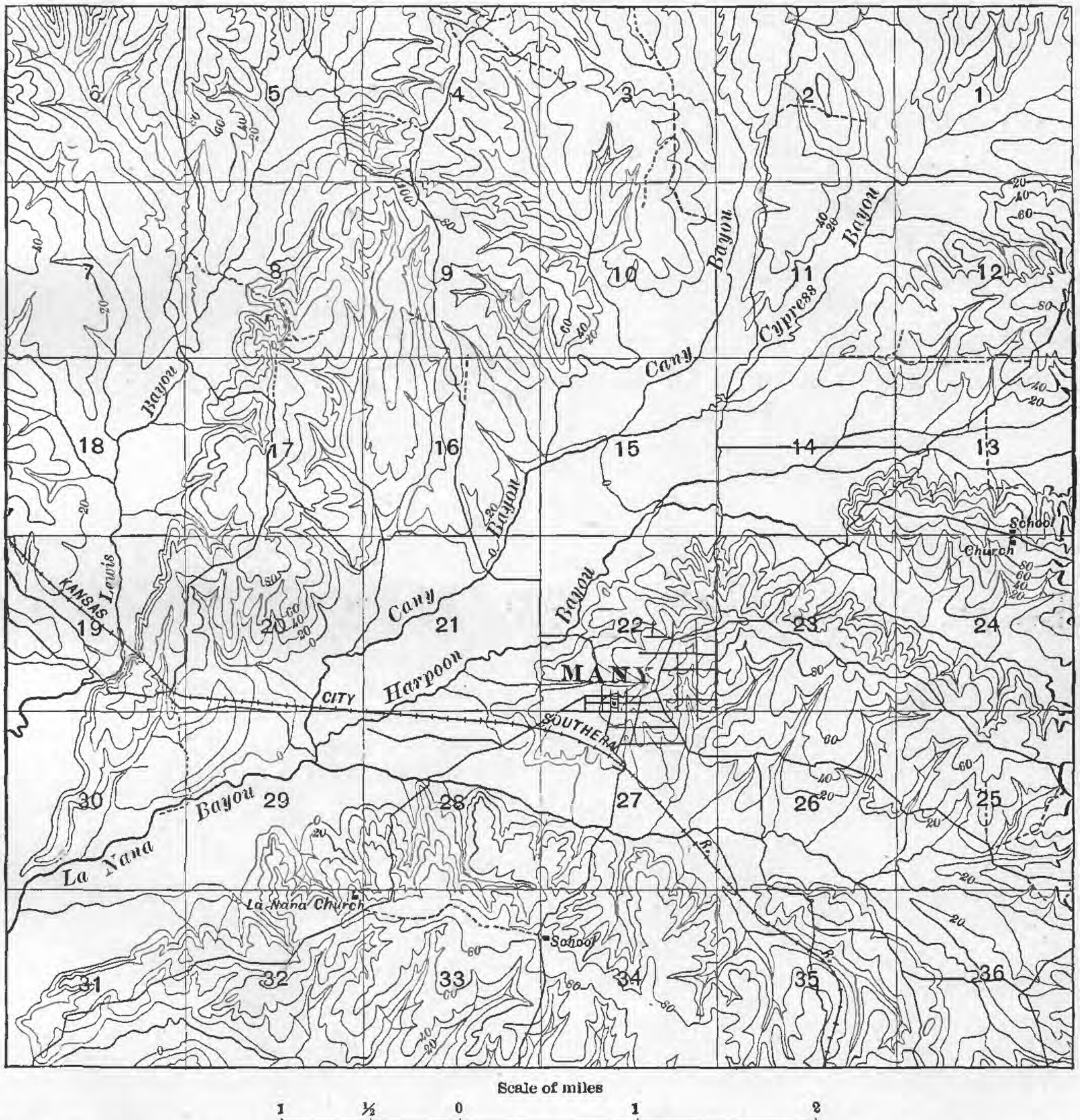


FIG. 7.—Sketch topographic map near Many, Sabine Parish, La., showing the characteristic flat-bottomed, steep-sided, small stream valleys of northern Louisiana and southern Arkansas, by A. C. Veatch, 1899.

Arkansas produced. Indeed, the only great difference between the topography of that day and this was that the principal valleys were 100 to 150 feet deeper, and the Port Hudson terraces (Pl. III, in pocket) were entirely absent. The valleys of the small streams did not then show their present anomalous, steep-sided, flat-

A. ON FLAT FORK CREEK, 3 MILES FROM TENEHA, TEX.

Showing tree-covered top and barren surrounding country. Substructure: Sabine Eocene.

B. ON FLAT FORK CREEK, 3 MILES FROM TENEHA, TEX.

Symmetrical form partially destroyed by stream erosion. Substructure: Sabine Eocene.

C. AT PRESCOTT, ARK.

Substructure; Arkadelphia Cretaceous.

D. ON PINE FLATS, NEAR TARBINGTON, TEX.

Substructure: Port Hudson (Quaternary). Photograph by Vernon Bailey.

NATURAL MOUNDS.

A

B

SMALL SAND CONES FORMING OVER GAS AND WATER VENT ON FLAT FORK CREEK, 3 MILES SOUTH OF TENEHA, SHELBY COUNTY, TEX.

C

MUD CONES NEAR DOUGLASTOWN, LONG ISLAND, N. Y.

Formed by pressure of underlying artesian water.

SAND AND MUD CONES.

ANT-HILL THEORY.

In considering the ant-hill hypothesis it must be conceded at the outset that in size and distribution these mounds exceed the work of any mound-building insects in this country. They are, however, approximated in size by some of the mounds of the leaf-cutting ants, the Atta. These are reported by Dr. W. M. Wheeler, formerly professor of zoology in the State University of Texas and an authority on ants, to reach a diameter in Texas of 40 to 50 feet and a height of 1 to 2 feet. He states that the hills are very stable and persist after the colony has migrated or become extinct. Mr. E. A. Schwarz, of the National Museum, reports that in Cuba the Atta hills often reach a height of 10 to 12 feet, with a diameter several times as great, and in places completely overrun the cane fields. These occurrences greatly reenforce the theory of an ant origin.

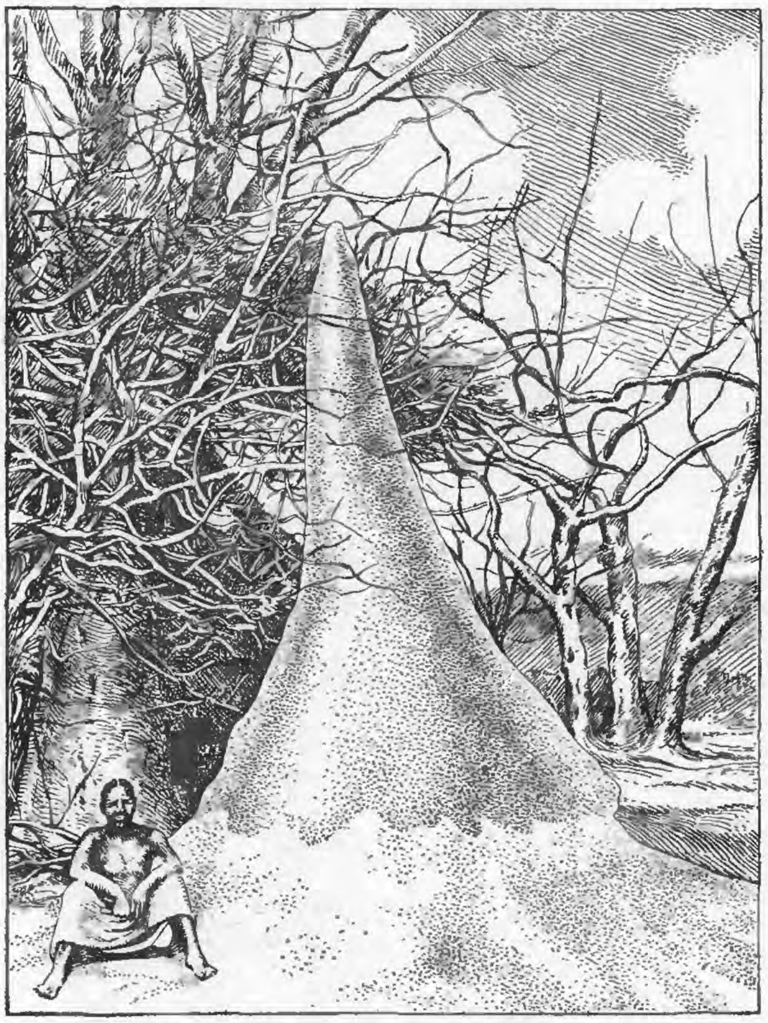


FIG. 12.—An African termite hill. (Drawing from photograph by Sir H. H. Johnston.[b]) Note the broad, low mound, on which the central spire rests, produced by wash of the central hill.

An alternative "ant theory" is that these mounds are the work of mound-building varieties of the so-called "white ants" (termites), which are notably developed in the tropical parts of South America and Africa and in Australia.[a] The immense hills of certain varieties of these termites, notably *Termes bellicosus*, which form a very important minor topographic feature over wide areas in Africa, are the nearest approach of any insect work to these natural mounds, both in size and bulk of material represented. These structures have a conical, sugar-loaf, or bee-hive shape and range from 6 to 20 feet in height and 50 or more feet in diameter (fig. 12). They are composed of mud in which more or less vegetable matter is mixed, and so, like the mounds, are best developed in clay regions. Should these cones be deserted by the termites, they would weather down into broad, low mounds which, because of their greater height and of the vegetable matter mixed with them, would have a looser character than the surrounding soil.

[a] For a discussion of termites see Encyclopædia Britannica, 10th ed., vol. 33, 1902, pp. 255-256, and the references there given.

[b] British Central Africa, New York, 1897, p. 371.

A. LOW CIRCULAR DUNES IN WHITE VALLEY, WESTERN UTAH.

Produced by sand or dust lodging about low desert vegetation. Photograph by G. K. Gilbert.

B. GOODWIN RAPIDS ON SABINE RIVER, NEAR COLUMBUS, LA.

One of the group of low shoals produced by recent upward movement along the Angelina-Caldwell flexure.

A. STEAMER "AID" REMOVING A PORTION OF THE GREAT RED RIVER RAFT IN 1873.

Note powerful lifting cranes and inclined platform bow, on which logs were hauled up for cutting with steam and hand saws. Location: 1 mile above Cedar Bluffs, La.; sec. 22, T. 21 N., R. 14 W. Photograph by R. B. Talfor.

B. ONE OF THE SEVERAL TIMBER JAMS COMPOSING THE GREAT RED RIVER RAFT.

In the region of slack water produced by these jams silt accumulated very rapidly and effectually filled the channel of the river (Pl. XXXI). Location: Channel of Red River, 5 miles below the Arkansas-Louisiana State line; sec. 29, T. 23 N., R. 14 W. Photograph by R. B. Talfor.

GREAT RED RIVER RAFT.

A. OPEN WATER AT FOOT OF ONE OF THE SEVERAL RAFTS WHICH MADE UP THE GREAT RED RIVER RAFT.

Timber jams and open water alternated throughout the length of the Great Raft, the open water being one-half to two-thirds of the whole. Location: One-half mile above head of Red Bayou, La., opposite Millers Bluff; sec. 16, T. 22 N., R. 14 W. Photograph by R. B. Talfor.

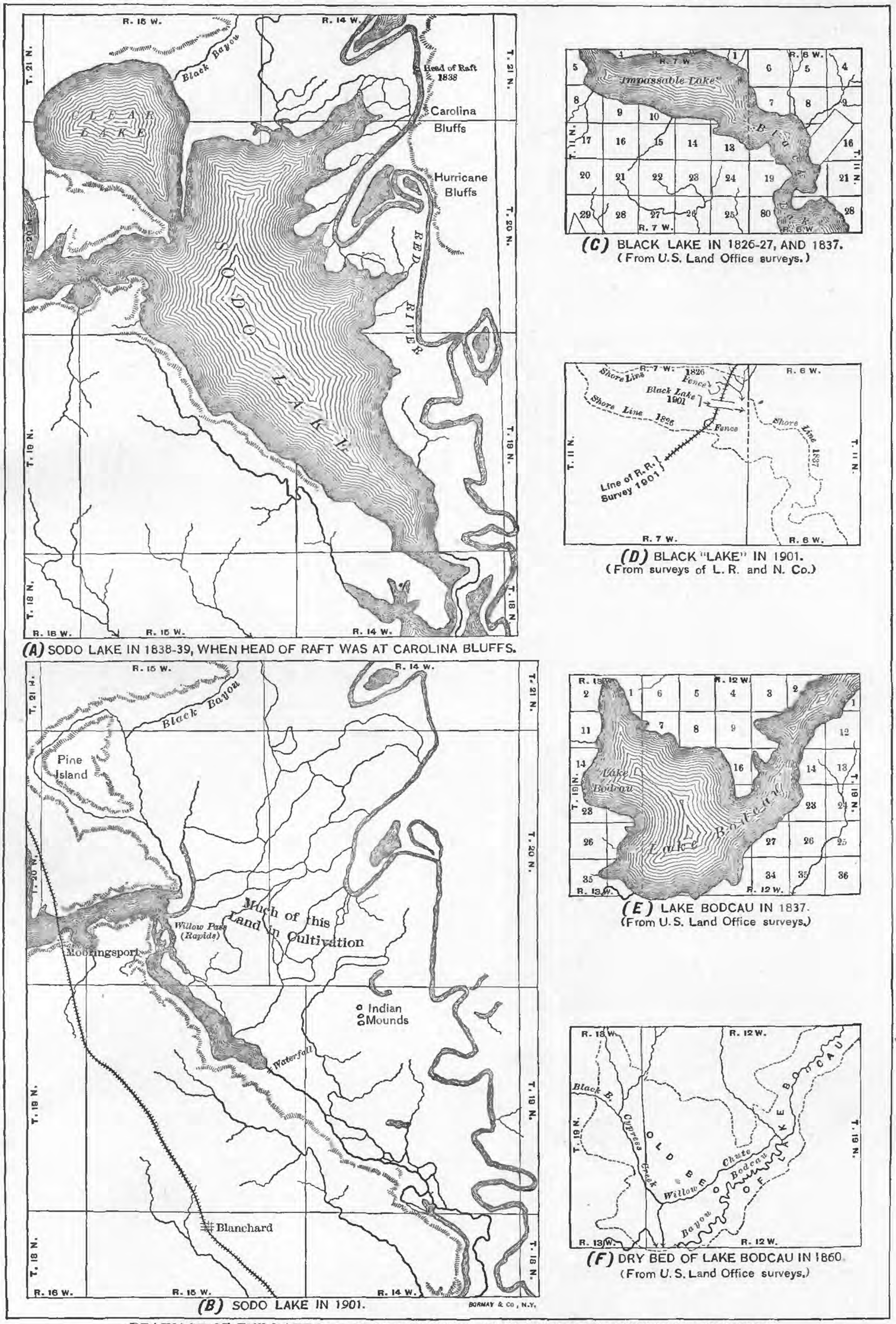
B. TIMBER DEADENED IN SHALLOW LAKE, PRODUCED BY THE RAFT.

Location: Sec. 33, T. 23 N., R. 14 W., Caddo Parish, La. Photograph by R. B. Talfor.

C. TIMBER DEADENED IN TEMPORARY RAFT LAKE WHICH WAS DRAINED BY THE REMOVAL OF THE RAFT.

Location: Head of McWillie Lake, sec. 16, T. 22 N., R. 14 W., Bossier Parish, La. Photograph by R. B. Talfor.

GREAT RED RIVER RAFT.



(A) SODO LAKE IN 1838-39, WHEN HEAD OF RAFT WAS AT CAROLINA BLUFFS.

(B) SODO LAKE IN 1901.

(C) BLACK LAKE IN 1826-27, AND 1837.
(From U.S. Land Office surveys.)

(D) BLACK "LAKE" IN 1901.
(From surveys of L. R. and N. Co.)

(E) LAKE BODCAU IN 1837.
(From U. S. Land Office surveys.)

(F) DRY BED OF LAKE BODCAU IN 1860.
(From U. S. Land Office surveys.)

DRAINAGE OF THE RAFT LAKES OF RED RIVER VALLEY IN NORTHWESTERN LOUISIANA.
BY A. C. VEATCH.

Scale
0 5 10 miles

ultimately the Mississippi. After the passage of the raft up the river this new course was so well established that it has remained the main channel of the river to this day.

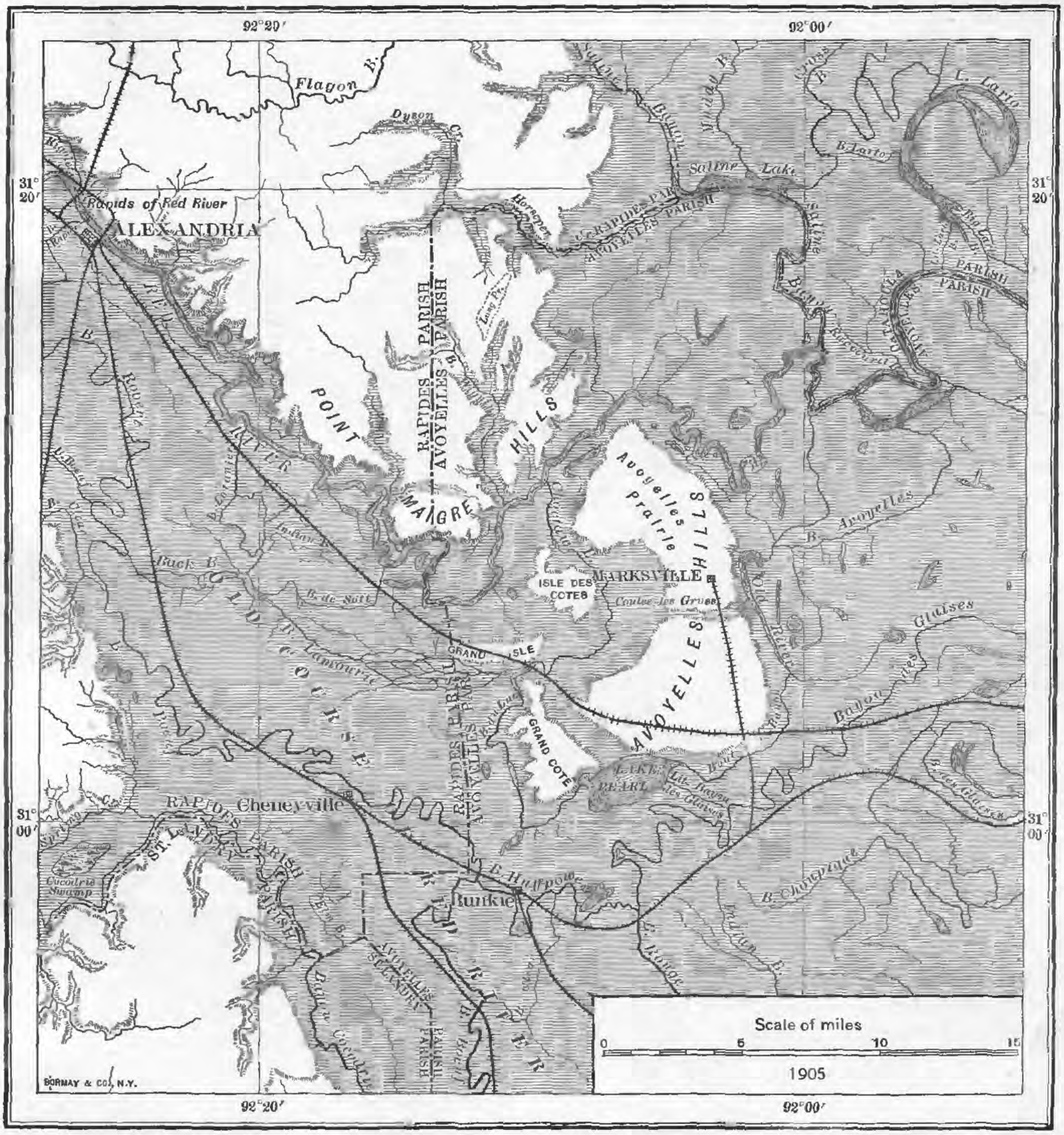


FIG. 13.—Map showing diversion of Red River below Alexandria, La., and location of the rapids. Shaded areas are land subject to overflow; white areas, terraces and hill lands.

PRODUCTION OF THE "RAPIDES" NEAR ALEXANDRIA, LA.

Near Alexandria the choking and building up of the old channel of Red River during the raft period and the free outlet afforded by the low place between the Avoyelles and Point Maigre hills caused the development of a new channel along the north side of the valley. When the river, in adjusting itself to its new course,

A. FERRY LAKE NEAR MOORINGSPORT, CADDO PARISH, LA.

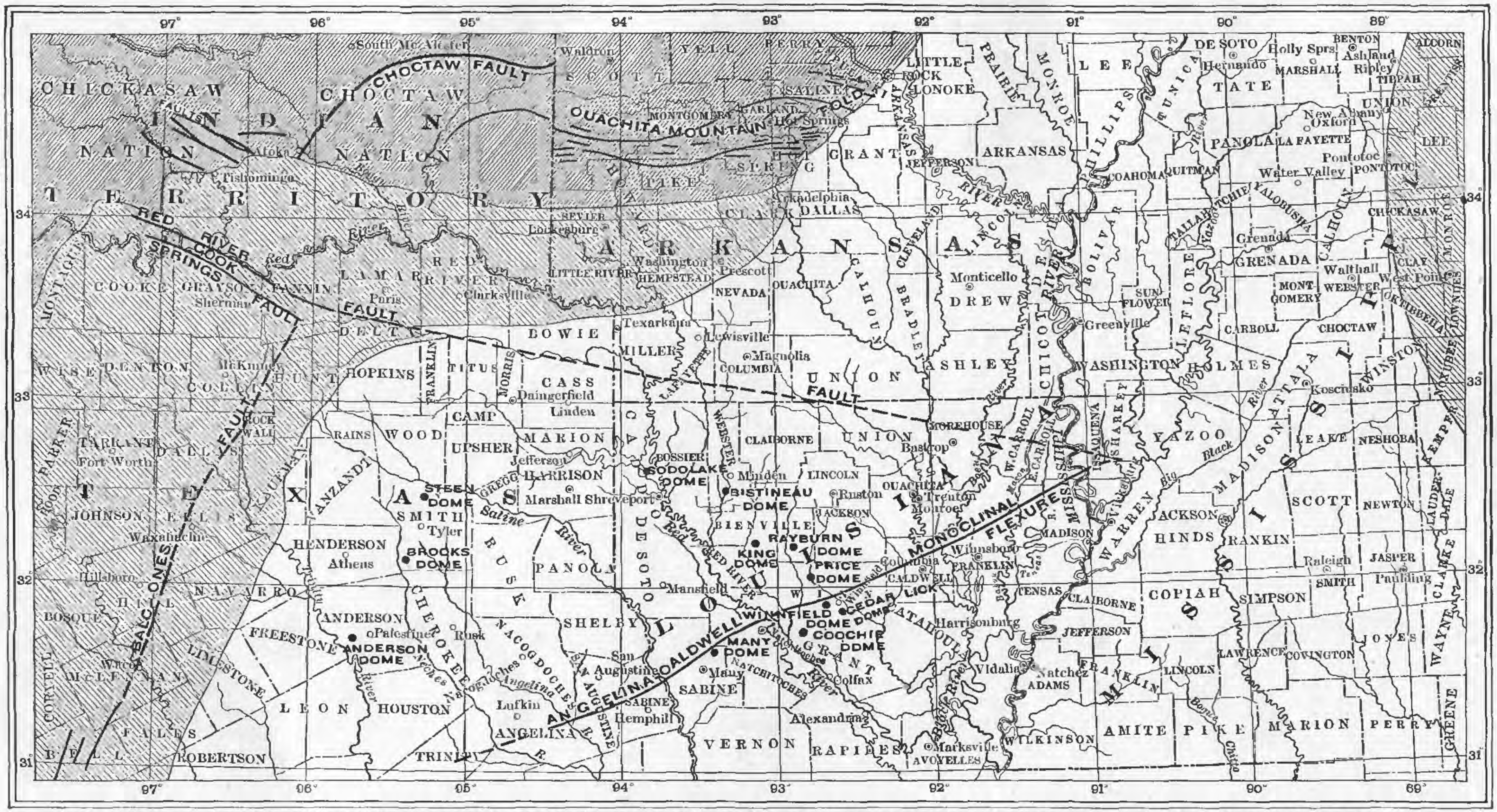
Showing stumps of pine, oak, and cypress killed by flooding during raft period, and now exposed by partial drainage of lake.

B. WAVE-FORMED BLUFF ON SOUTH SHORE OF FERRY LAKE.

Partial drainage of lake by removal of raft has left the bluff 15 feet above water level. Note trees growing on old beach.

FERRY LAKE, A PARTIALLY DRAINED RAFT LAKE.

1393—No. 46—06——6



PRINCIPAL STRUCTURAL FEATURES OF THE COASTAL PLAIN IN NORTHERN LOUISIANA AND SOUTHERN ARKANSAS WITH SOME OF THE STRUCTURAL FEATURES OF ADJOINING AREAS.

Scale

40 0 40 80 120 miles

1905

LEGEND

Tertiary with Quarternary Covering in places.

Cretaceous.

Archean and Paleozoic.

or impervious material, and spraying its center with water. When the lower portion of the pile has become saturated, little streams will begin to trickle from one or more points at the base, the number depending on the shape of the ground where the pile is situated. These little rivulets represent springs and it will be found that the water will flow for some time after the cessation of the spraying, the length of

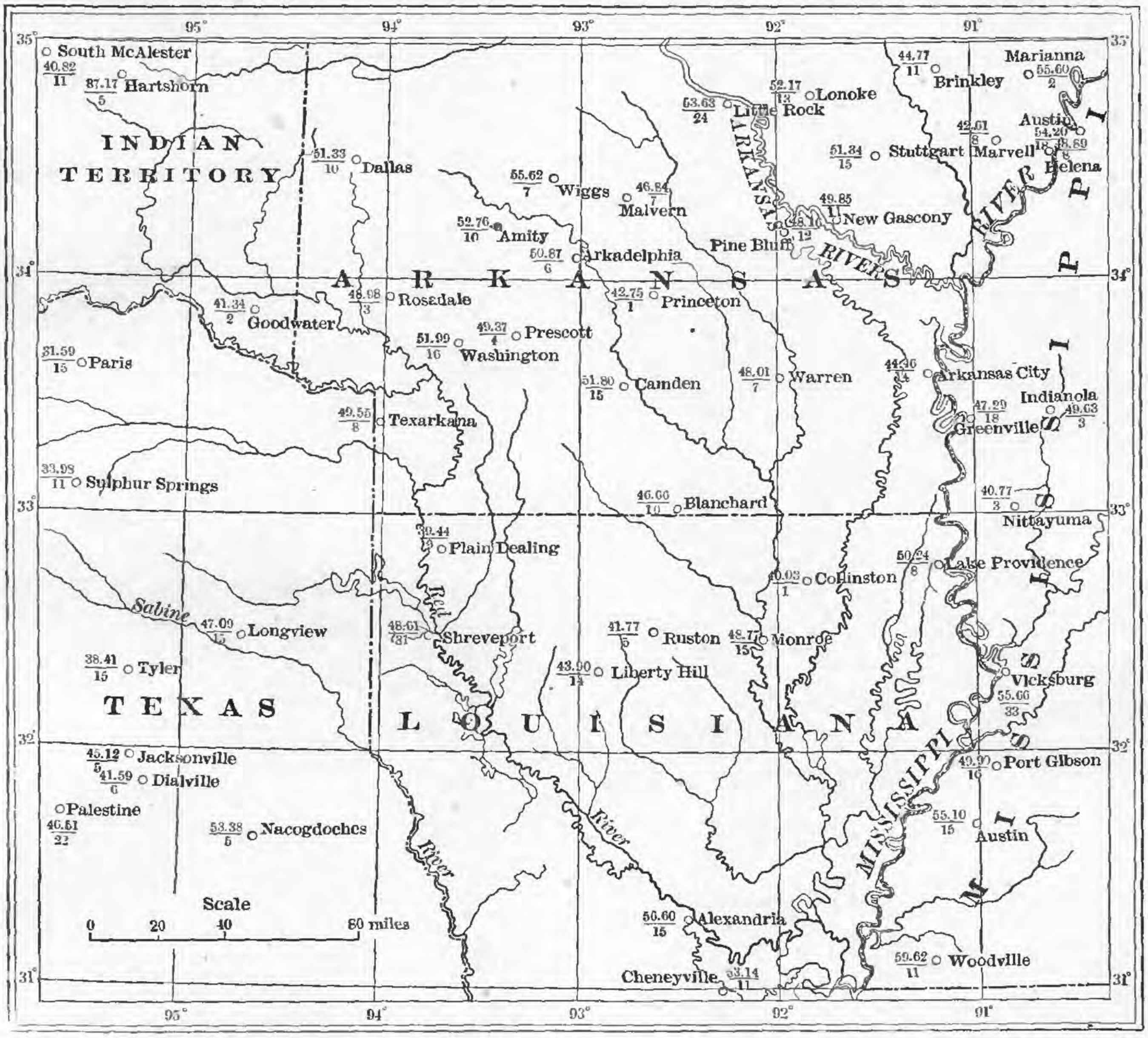
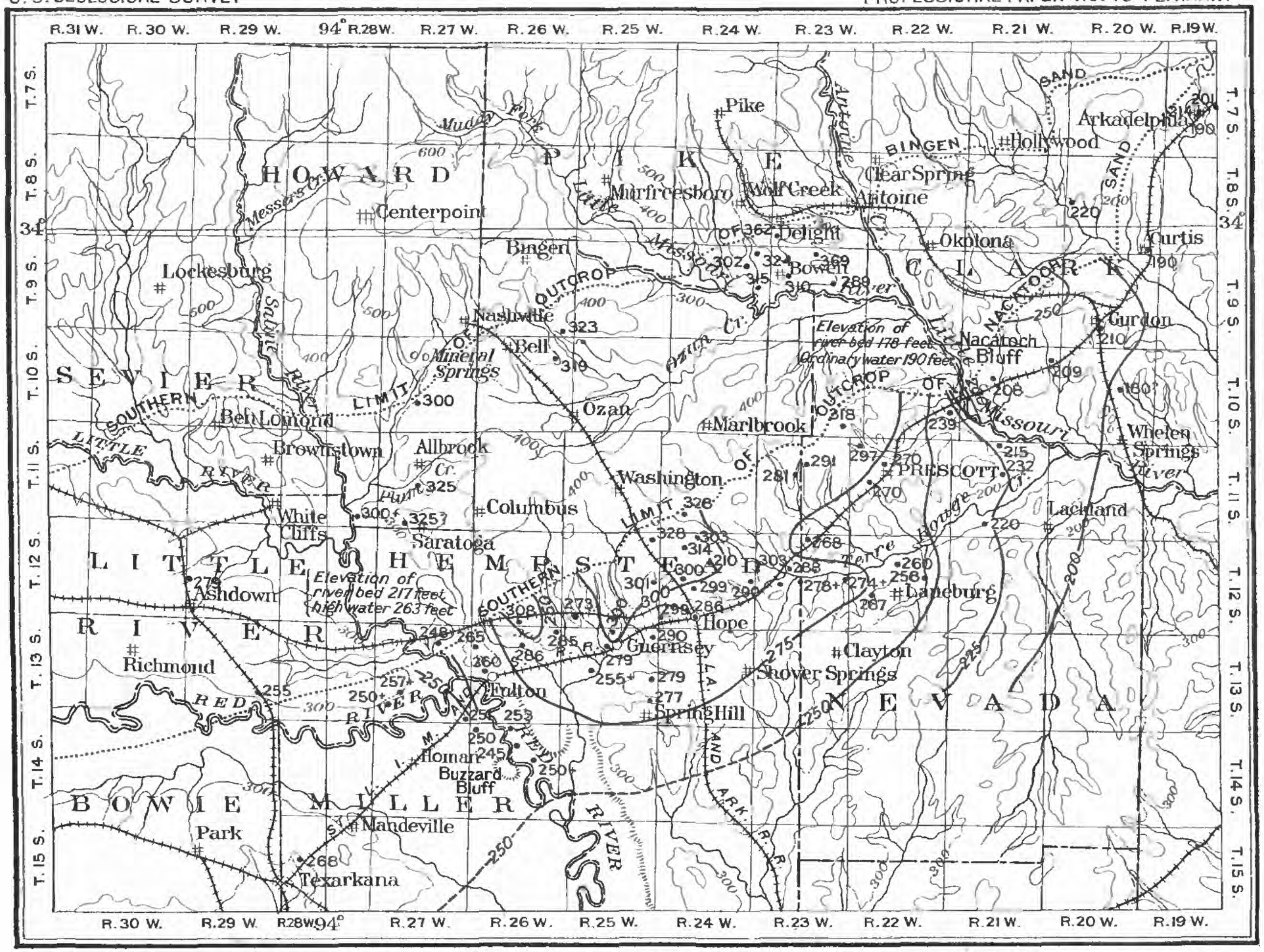


FIG. 18.—Annual rainfall in northern Louisiana and southern Arkansas, compiled from reports of the United States Weather Bureau for the years 1902 and 1903. Lower figures represent period of years; upper figures, average rainfall for period given.

time depending on the size of the pile and the coarseness of the particles composing it. Imagine this sand pile increased to many feet in thickness and covering the top of a hill or group of hills which is underlain by clay beds, and an idea will be had of the character and cause of many of the springs in northern Louisiana and southern Arkansas.[a]

[a] For another type of springs see p. 76.



LEGEND

•325 Water level in feet above sea level in Bingen sand wells

•286 Water level in feet above sea level in Nacatoch sand wells

250 Lines of equal head in Nacatoch sand

MAP OF SOUTHERN ARKANSAS SHOWING VARIATIONS IN HEAD OF WATER IN NACATOCH AND UPPER BINGEN SANDS

By A.C.Veatch

Surveyed in 1902

Scale

0 10 20 30 miles

CONTOUR INTERVAL 100 FEET

1905

JULIUS BIEN & CO.N.Y.

feet at Boughton (fig. 23). The decrease of pressure and the rate of fall of the creek are not exactly the same, and the irregular variation of these two curves gives rise to several disconnected artesian areas.

Wells near the points where the outcrop of the Nacatoch sand crosses the larger rivers fluctuate with the stage of the river. (See well 203.) In some cases this is due to the fact that the height of the point of discharge for the water moving laterally from the high points in the outcrop along the divide changes with any

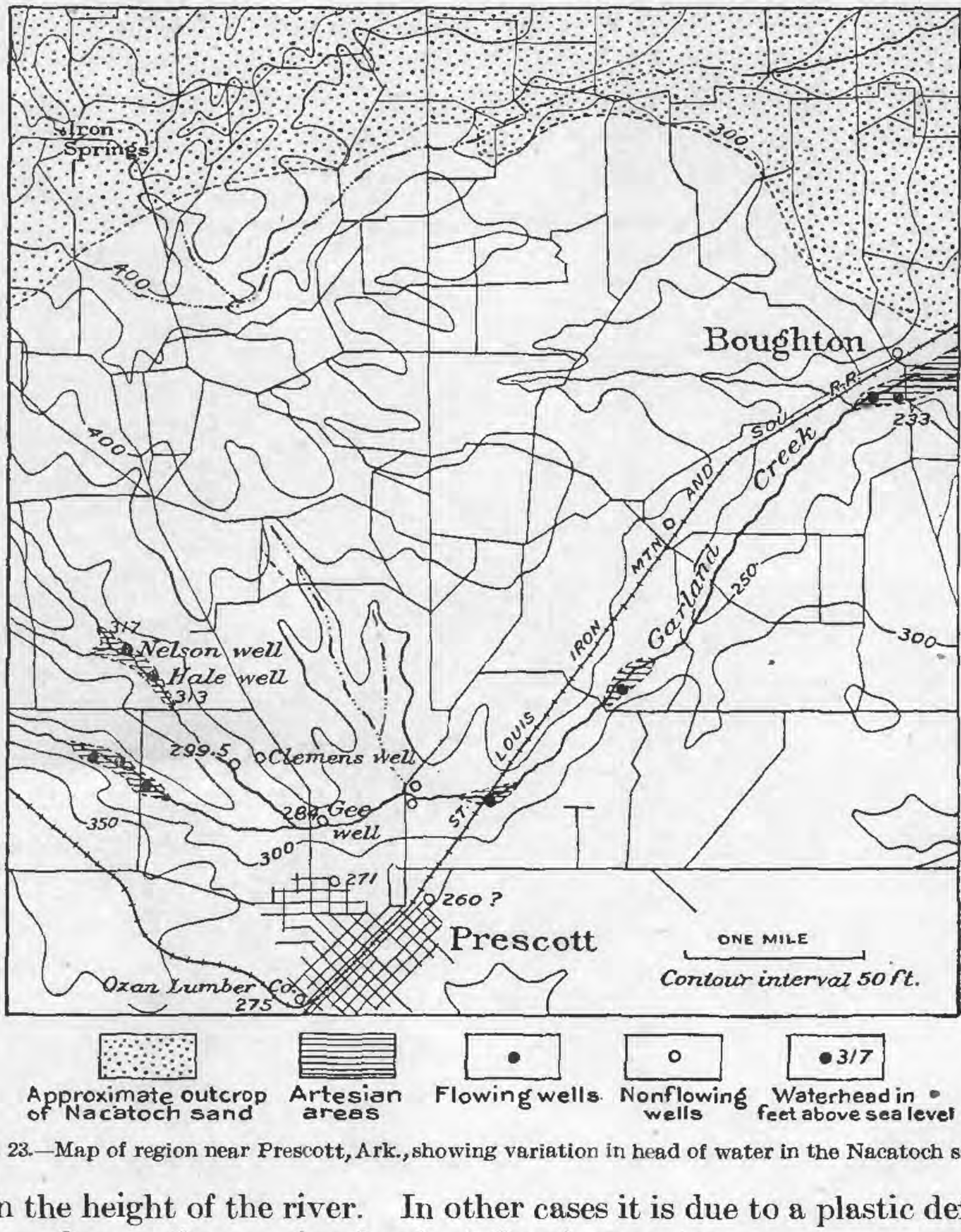
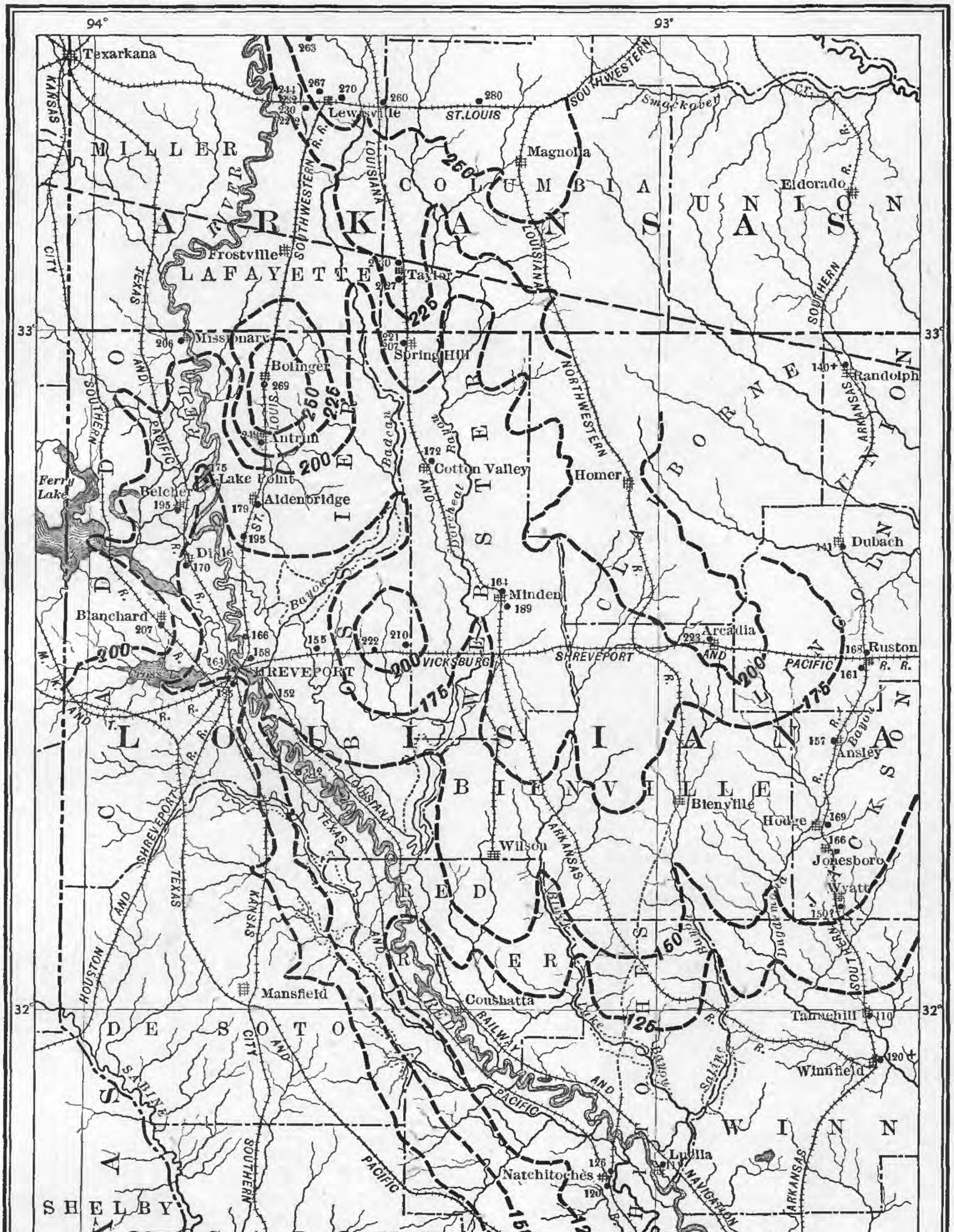


FIG. 23.—Map of region near Prescott, Ark., showing variation in head of water in the Nacatoch sand.

variation in the height of the river. In other cases it is due to a plastic deformation of the strata, due to the varying head of the river waters.

Quality.—The Nacatoch sand, like the sub-Clarksville sand, yields a soft alkaline water whose total mineral content varies considerably from place to place. Where allowed to stand in contact with the overlying blue clays, as it does in the majority of nonflowing domestic wells, which are cased only through the layer of "caving dirt," generally 30 or 40 feet thick, it becomes quite hard; but in freely flowing wells and in wells which are cased to the bottom it is soft. It has

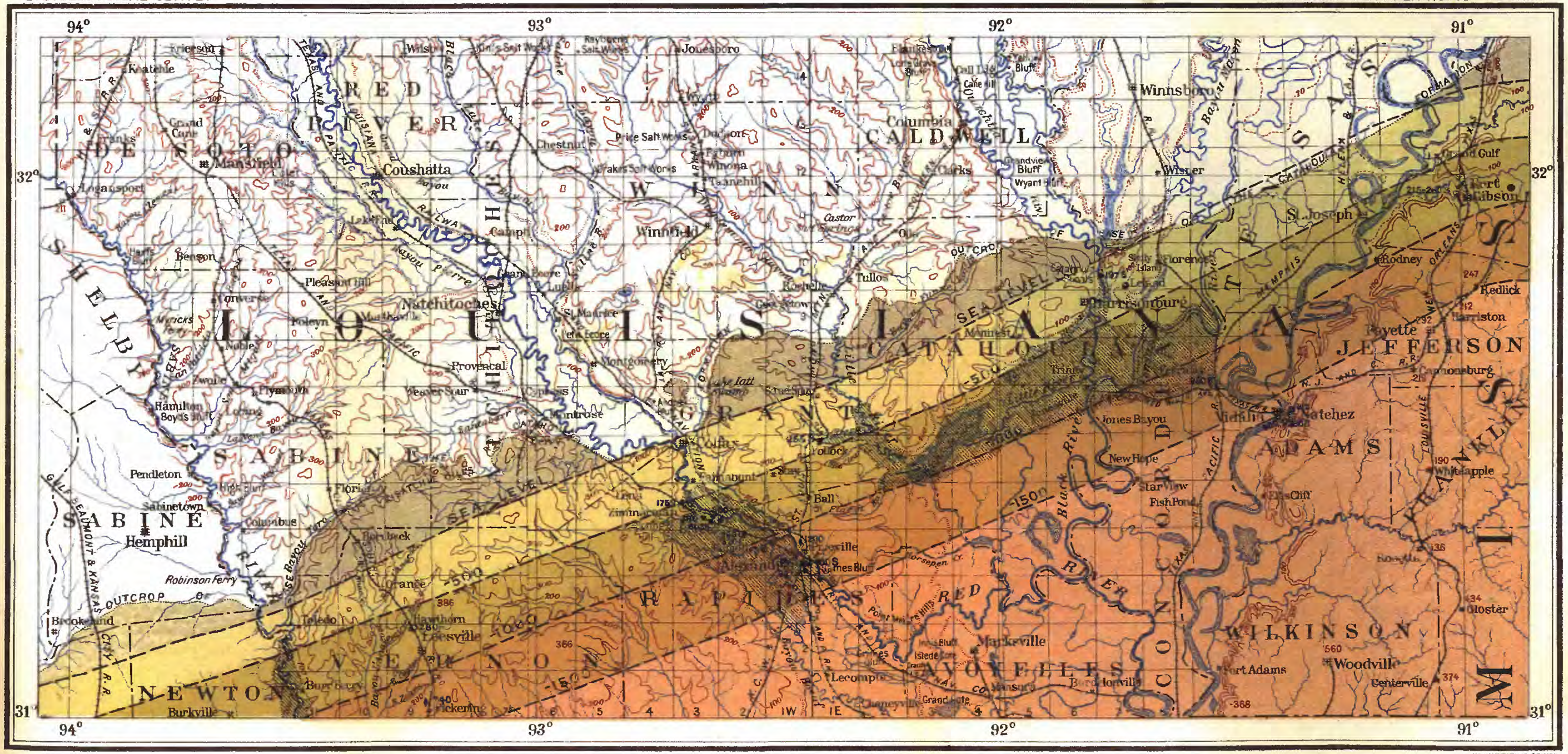


●191 Elevation to which water will rise, in feet above sea level.

200 Approximate lines of equal head.

MAP SHOWING VARIATION IN HEAD OF WATER IN THE SABINE SANDS IN NORTHWESTERN LOUISIANA AND SOUTHWESTERN ARKANSAS.

By A. C. Veatch, 1903.



LEGEND

- Area not underlain by Catahoula sands

DEPTHS OF BASE OF CATAHOULA FORMATION

- Above sea level
- 0-500 feet below sea level
- 500-1000 feet below sea level
- 1000-1500 feet below sea level
- Over 1500 feet below sea level

- Area of flowing wells
- ● Flowing wells
- ○ Nonflowing wells
- ♦ Failures

Figures give depths from the surface, in feet, of water-bearing strata or bottom of wells

Letters give quality of water
B Brine
S Soft

Note: Depths given by shading and contours refer to basal layers. Wells can generally be completed at less depths. In the area underlain by the Catahoula formations it is not advisable to continue wells below the depths indicated. See Chapter IV for a discussion of conditions in each locality

JULIUS BIEN & CO. N.Y.

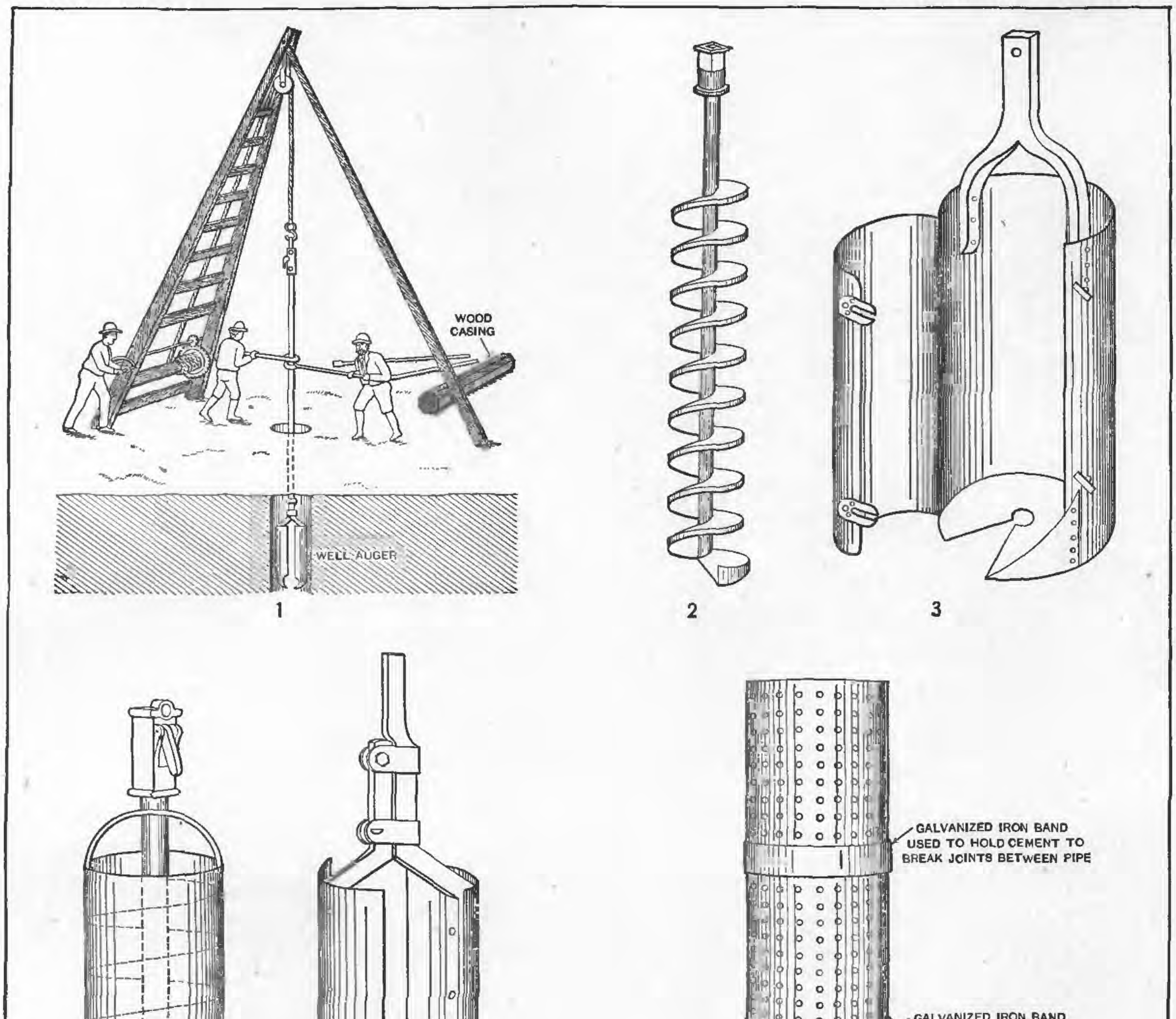
MAP OF THE CATAHOULA ARTESIAN RESERVOIRS IN CENTRAL LOUISIANA
BY A.C. VEATCH

Surveyed in 1902

Scale

10 5 0 10 20 30 40 50 miles

1905



WELL-BORING AND WELL-PUNCHING TOOLS.

1. Ordinary well-boring outfit.
2, 3, 4, 5. Well-boring augers.
6. Self-fastening couplings used on Challenge well-auger poles.
7. 18-inch perforated straight tile, used as a strainer in Elliott tile wells. Note also Elliott patent iron shoe.
8. Well punch.

WELL MAKING.

TOOLS OF AN ARKANSAS WELL-BORING OUTFIT.

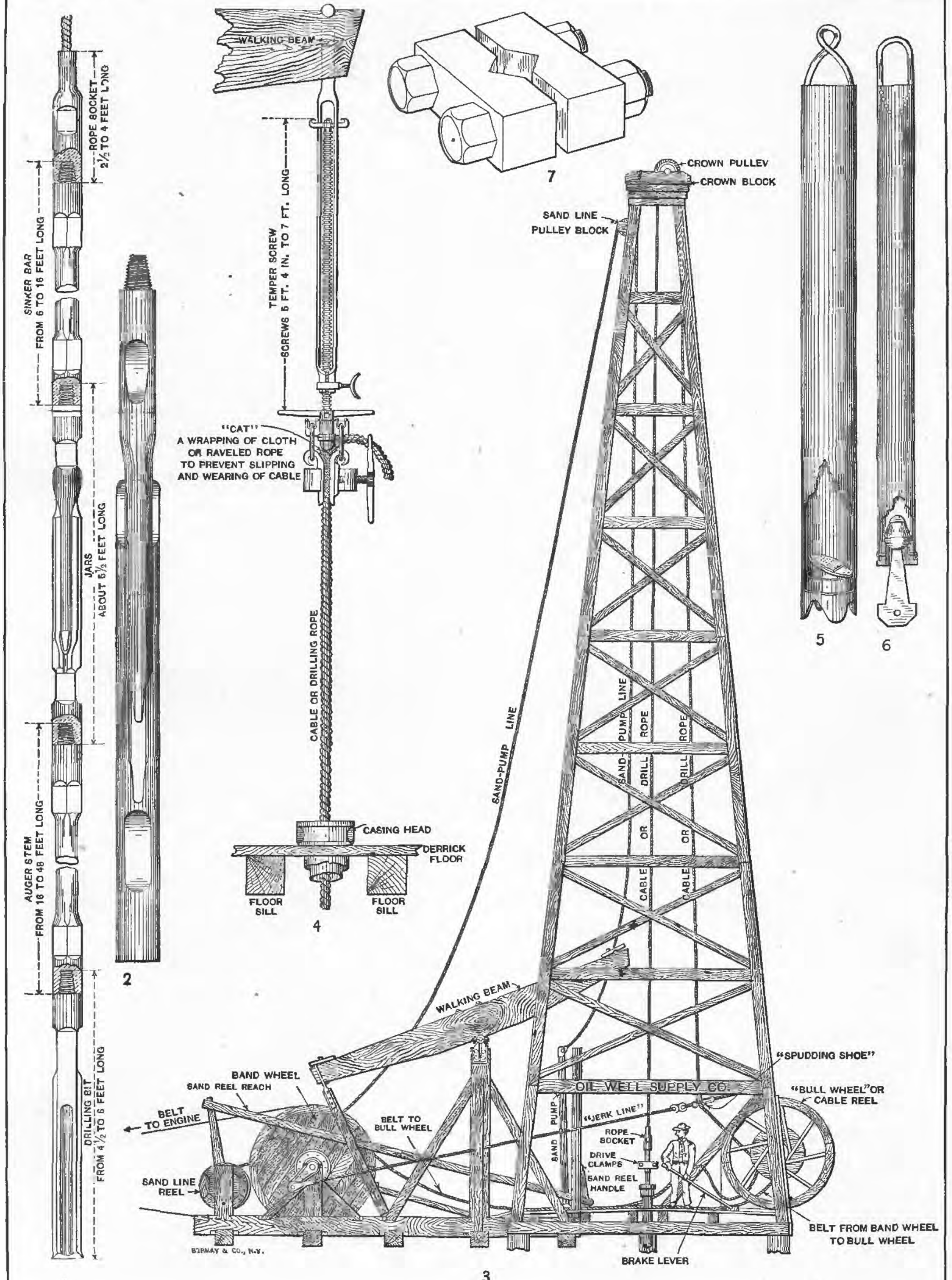
WELL MAKING.

A

B

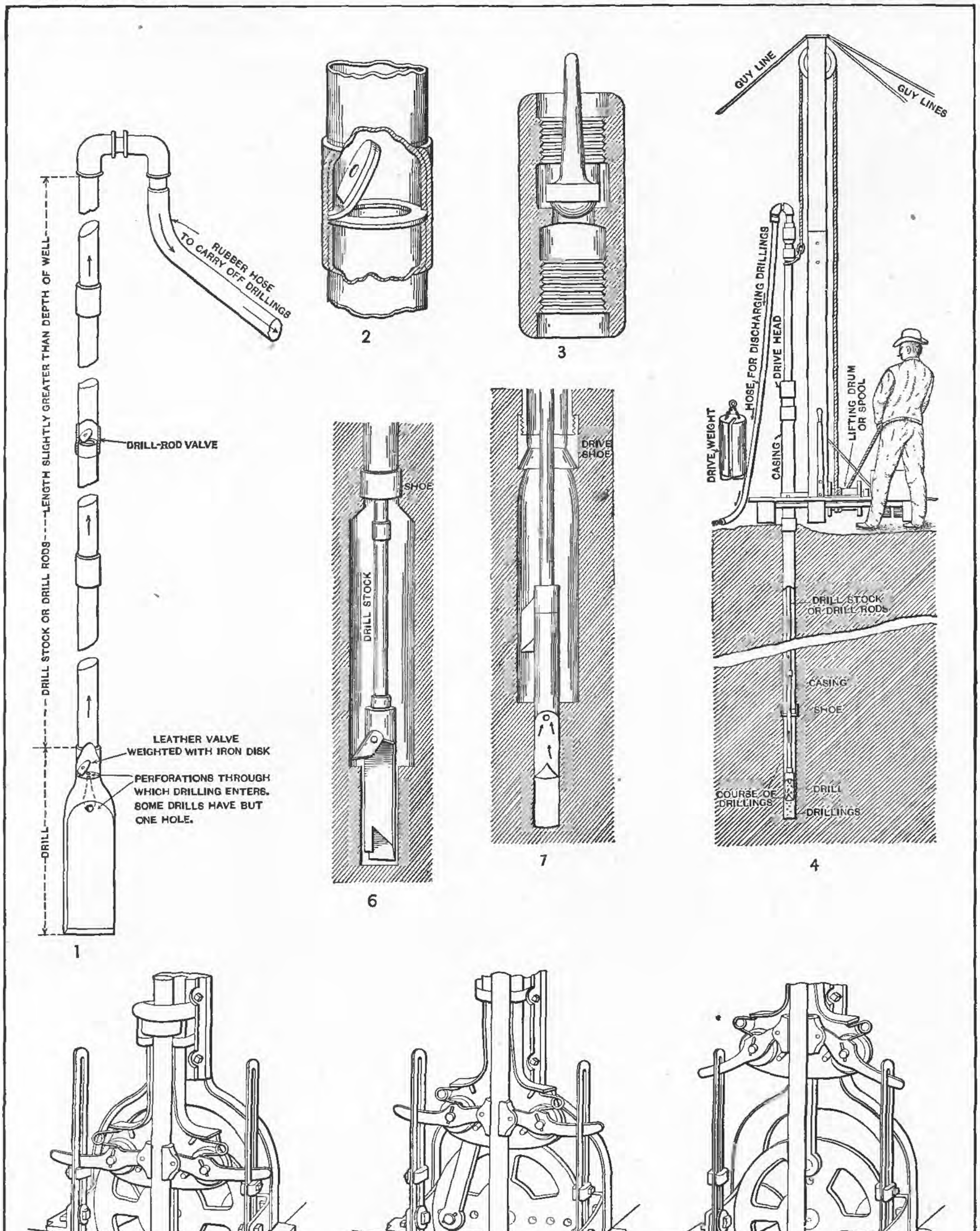
WELL-BORING OUTFIT OF MR. G. B. HIPP, OF GARLANDVILLE, ARK.

WELL MAKING.



CABLE RIG OR DROP-DRILL OUTFIT.

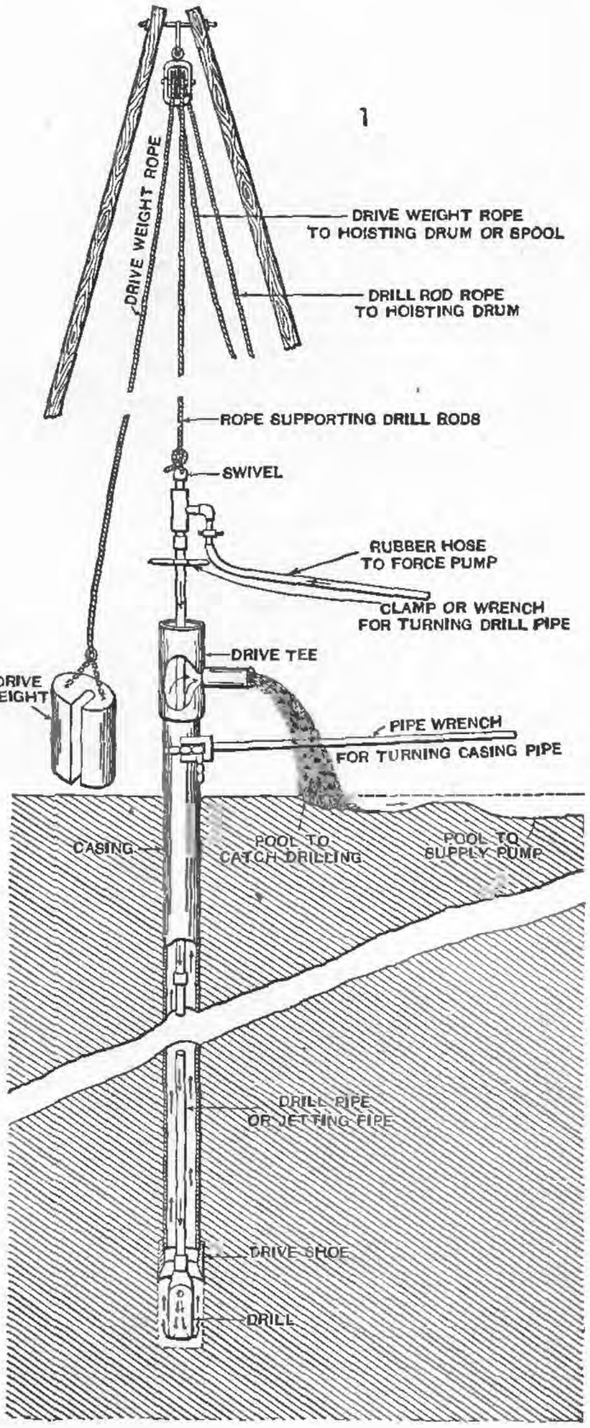
WELL MAKING.

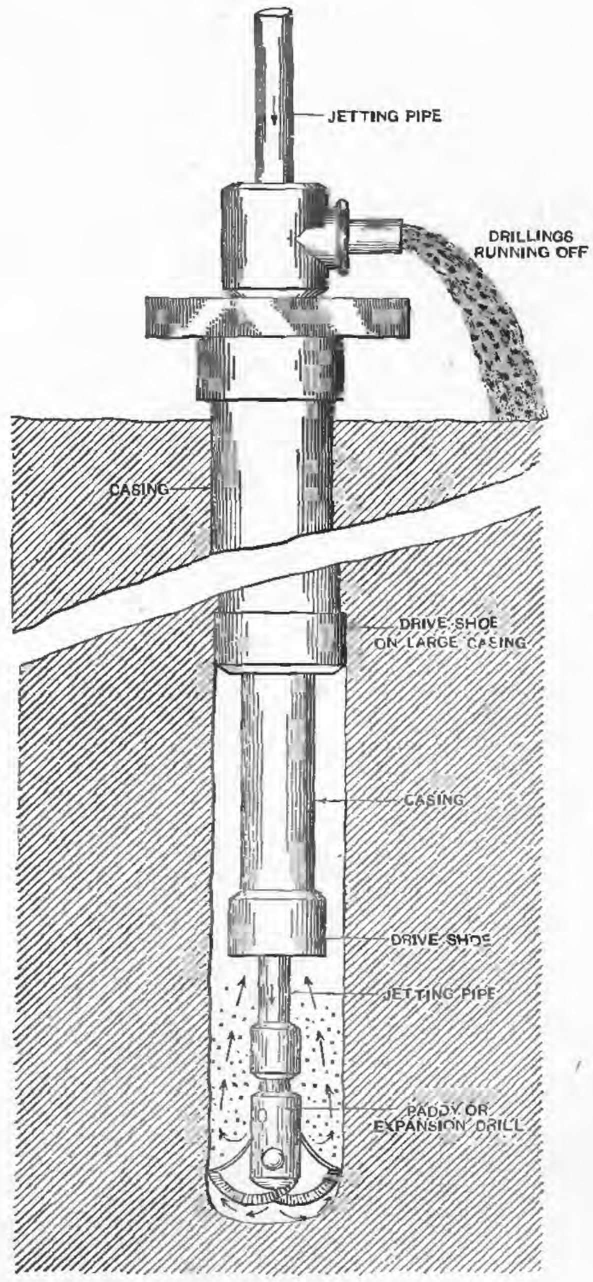


AUTOMATIC SAND-PUMPING PROCESS.

WELL MAKING.

1





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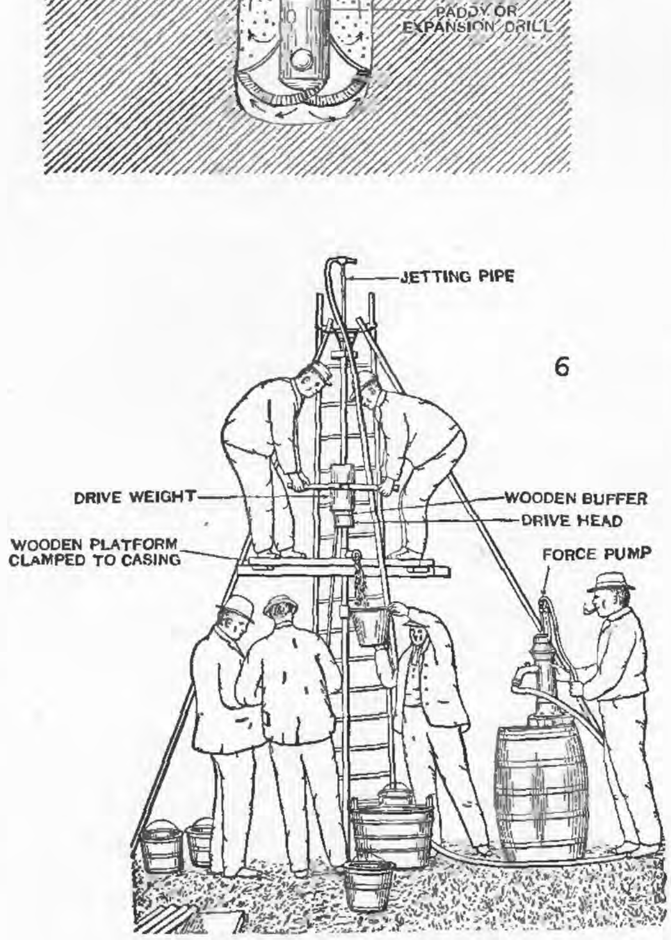
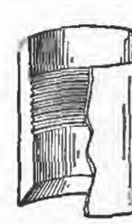


3

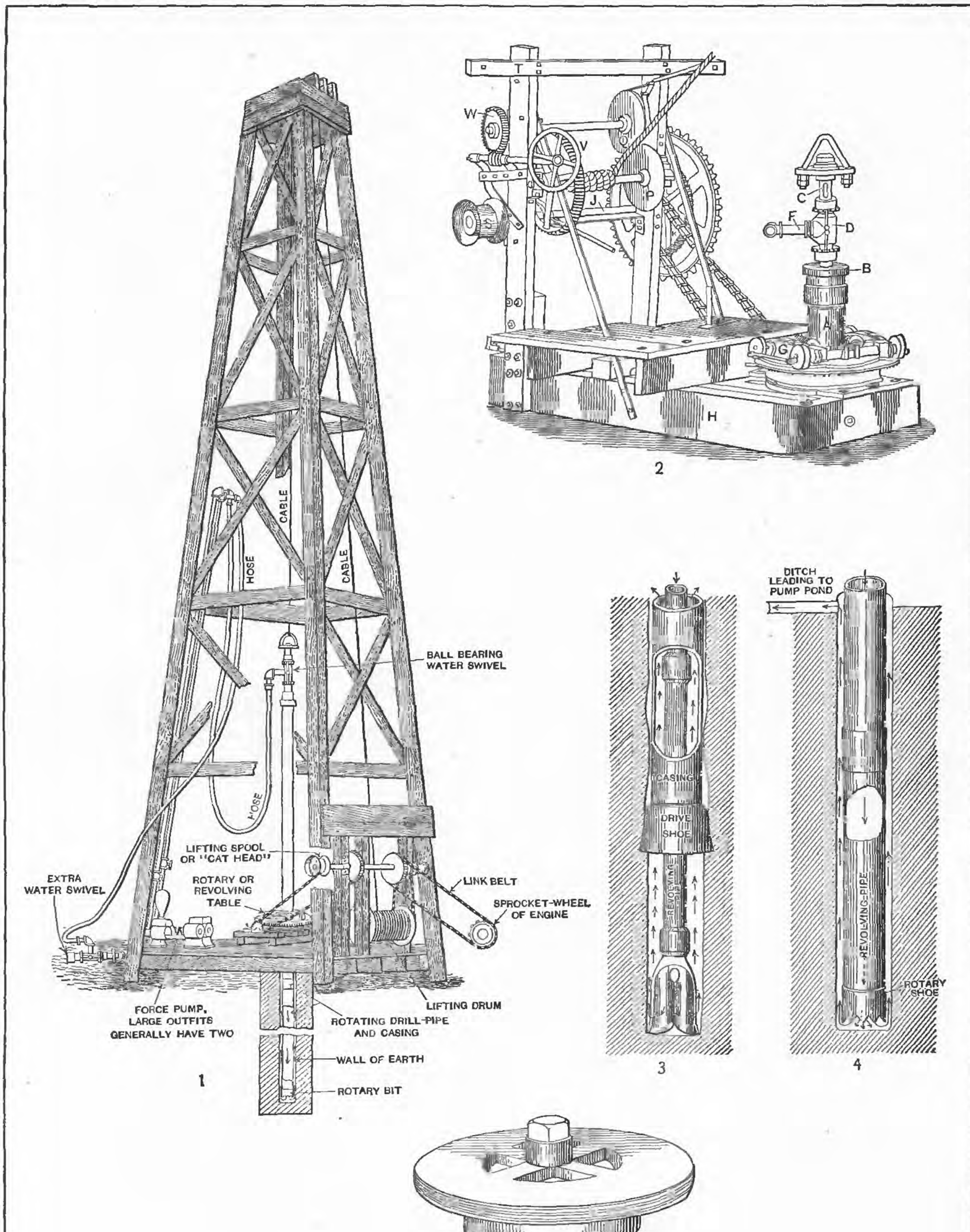
4



5

JETTING PROCESS.

WELL MAKING.



ROTARY PROCESS.

WELL MAKING.