The Effect of the Addition of Heat from a Powerplant on the Thermal Structure and Evaporation of Lake Colorado City, Texas *By* G. EARL HARBECK, JR., G. E. KOBERG, *and* G. H. HUGHES STUDIES OF EVAPORATION

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PREFACE

This report was prepared in the Water Resources Division of the Geological Survey under the direct administrative supervision of R. W. Davenport, Chief, Technical Coordination Branch, followed by C. C. McDonald, Chief, General Hydrology Branch, and under the technical guidance of W. B. Langbein. Most of the fieldwork and data processing was performed by personnel of the Surface Water Branch, J. V. B. Wells, Chief, under the supervision of Trigg Twitchell, District Engineer. The study was conducted in cooperation with the Board of Water Engineers, State of Texas, H. A. Beckwith, Chairman, succeeded by R. M. Dixon.

The cooperation of the Texas Electric Service Co., Fort Worth, Tex., who assisted with the installation of equipment and whose personnel made many of the routine daily observations, is greatly appreciated.

The help of the U. S. Weather Bureau in installing a pan evaporation station at the reservoir is gratefully acknowledged. These data have been furnished to the Weather Bureau for analysis and the results are to be published in a separate report.

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STUDIES OF EVAPORATION

THE EFFECT OF THE ADDITION OF HEAT FROM A POWERPLANT ON THE THERMAL STRUCTURE AND EVAPORATION OF LAKE COLORADO CITY, TEXAS

By G. EARL HARBECK, JR., G. E. KOBERG, and G. H. HUGHES

ABSTRACT

Lake Colorado City, a reservoir in north-central Texas, is used as a source of cooling water for a thermal-electric powernlant. Evaporation from the lake was determined by the energy-budget method for the period July 1954 to October 1955. Annual evaporation from Lake Colorado City was 93 inches, of 'vhich 85 inches was natural evaporation and 8 inches resulted 'Tom addition of heat to the lake by the powerplant.

Analyses indicate that if no heat had been added by the nowerplant, the water-surface temperature would have been only 0.8° C lower than that observed. The temperature rise is almost directly proportional to the heat input, so that doubling 'he quantity of heat added by the powerplant would raise the ' emperature an additional 0.8° C, if the volume of water in the "eservoir was the same as in 1954-55. The temperature rise in ". nearly empty reservoir, of course, would be much greater, and "n a full reservoir, less.

The increase in evaporation from Lake Colorado City, when expressed as a volume, is directly proportional to the amount of heat added and is practically independent of reservoir con- 'ents. During 1954-55 the volume of forced evaporation rerulting from the heat added by the powerplant, 910 acre-feet, was almost exactly the same as the volume of water diverted to Colorado City for municipal purposes.

The entire lake is being effectively utilized in disposing of heat, Water temperatures in the lower basin of the lake were higher :: n winter than those in the upper basin of the lake; no appreciable r'ifferences were observed in summer. Density differences be- ''ween the two parts of the lake were very small at all times.

A comparison between average lake temperatures (as determined from the thermal surveys) and plant intake water temperatures indicates that water is withdrawn from all levels of the lake above the pump intakes. The average withdrawal temperature to be expected, if the amount of heat added by the powerplant is increased, probably will be about equal to the rnticipated surface temperature, which can be determined from graphs in the text.

INTRODUCTION

Many thermal electric powerplants in the United Ftates withdraw water from a stream or lake, use it to condense steam, and return it to the stream or lake, practically undiminished in quantity, but with its temperature substantially raised. Obviously, evapora-

tion from the water body is increased. This usually is of little consequence where water supplies are ample, but it may be of considerable importance where supplies are short, as in the arid West.

If ample supplies of water are available from a take, perennial stream, or well, the decision whether to use cooling towers would be based only upon economy of construction and operation of the cooling system. If the amount of water consumed must be minimized, a reservoir or cooling pond will use less water than a cooling tower. Where natural streamflow is not adequate at all times, storage must be provided.

The design of cooling ponds has long been based on empirical formulas and on the judgment and experience of the designers. The prediction of water temperatures in a lake or cooling pond has been largely educated guesswork. The purpose of this report is to devise ways to predict the increase in evaporation and in water temperature of a reservoir when it is used as a cooling pond to dispose of given amounts of heat.

Representatives of the Texas Board of Water Engineers and the Geological Survey agreed that Lake Colorado City in north-central Texas would be suits ble for a field study. This reservoir, which is owned by the Texas Electric Service Co. of Fort Worth, seemed to have many advantages as a site for the study. For example, measurements of water volumes and temperatures could be made with adequate accuracy, and employees of the power company were available and eager to cooperate in making many of the routine deliy observations.

DESCRIPTION OF THE RESERVOIR AND DAM

Lake Colorado City is situated on Morgan Creek, a small tributary to the Colorado River in north-central Texas (fig. 3). The dam was built in 1949 by the Texas Electric Service Co. to provide a supply of cooling water for a steam-electric powerplant constructed

i.t the same time on the shore of the lake. The reser- "•oir area, is owned by the power company, but recreational use of the upper half of the lake is encouraged. Access to the lower half of the lake is restricted.

The capacity of the reservoir at service spillway level •;s 31,800 acre-feet, and the corresponding surface area •:s 2,030 acres. Since its construction in 1949, the •"eservoir has not spilled. Water is pumped from the •eservoir to Colorado City for municipal use.

The dam is a rolled-earth structure approximately 4,800 feet long. A morning glory spillway near the "'outhwest end of the dam is designed to handle most ?ood discharges For larger floods there is a long emer- •rency spillway between the plant and the dam that is designed to handle floods above the capacity of the Tiorning glory spillway.

Water is taken from Lake Colorado City at the power- •^lant, pumped to the plant where it is used for cooling, and returned to the lake through a canal. The canal discharges into the reservoir at a point nearly a mile distant from the plant. Water flows over a weir and "alls freely into the reservoir.

The drainage area of Morgan Creek above the dam "3 267 square miles. Average inflow for the period 1948-55, as determined from records obtained at a "ormer gaging station on Morgan Creek and from records of change in contents of the reservoir, was estimated to be 17.3 cubic feet per second (cfs). Average :nflow during the period July 1954 to October 1955 was 13.4 cfs. For periods as long as several months, owever, inflow is practically nil, so that natural streamflow could not have been relied upon as a source of cooling water, and it was necessary that storage be provided.

CLIMATOLOGY

According to Thornthwaite (1948) the climate of the Colorado City area is semiarid. The average annual temperature at San Angelo is 66.2° F, normal rainfall is 19.8 inches, and average wind speed is 10.4 miles per vour. San Angelo, the nearest Weather Bureau station. •3 approximately 75 miles south-southeast of Lake Colorado City. Approximately one-fourth of the annual rainfall occurs in April and May, another one- "ourth in September and October, with the remainder "airly well distributed among the other 8 months.

The data obtained at Lake Colorado City during the •^eriod July 1954 to October 1955, provide a basis for comparison between climatic conditions at Lake Colorado City and San Angelo. There is no appreciable difference between air temperature at the two 'ocations (fig. 4). Wind speed at Lake Colorado City is approximately 75 percent of that measured at San .Angelo (fig. 5), but the anemometer at San Angelo is **484263—59———2**

63 feet above the ground, as compared to approximately 40 feet above the water surface at Lake Colorado City, which could account for the difference. The correlation between vapor pressures at the two locations is poor (fig. 6), but there appears to be no significant bias.

During the period July 1954 to October 1955, the air temperature at San Angelo was generally slightly above normal, as was wind speed. Wind direction was not measured at Lake Colorado City, but during the same period monthly prevailing wind direction at Abilene, approximately 75 miles east of the lake, was south to south-southeast, and at San Angelo south to west-southwest. Humidity was slightly below normal.

Although rainfall at San Angelo during this period was only slightly more than half of normal, rainfall at Abilene was nearly normal. From these data it was concluded that weather conditions at Lake Colorado City were not greatly different from normal and that the period was reasonably representative of average conditions.

INSTRUMENTATION

WATER BUDGET

Although the evaporation from Lake Colorado City was computed by the energy-budget method it was necessary to obtain a reasonably accurate water budget to evaluate the advection and storage terms of the energy budget, and as an approximate check to guard against gross errors in the computed figures of evaporation.

INFLOW

Morgan Creek is the principal source of surface inflow to Lake Colorado City. Records of inflow are available at two stream-gaging stations, Morgan Creek near Westbrook, Tex., and Graze Creek rear Westbrook, Tex. The stations are located about 10 miles upstream from the bridge on U. S. Highway 80 which crosses the upper end of the lake (fig. 3). Standard stream-gaging techniques, as described by Corbett (1943), were used. Streamflow records for the two stations were classified as good, which means that the error in daily records is believed to be less than 10 percent.

During the 468 days of observation, surface flow occurred on 71 days; flows of 2 acre-feet or more per day occurred on 52 days. Figure 7 shows the monthly in-

flow to Lake Colorado City during the period of observation. During 5 periods of from 2 to 5 days measured inflow was more than 700 acre-feet. Comparison with the change in contents at Lake Colorado City during these periods indicates that some of the measured inflow apparently did not reach the reservoir. This will be discussed later.

Surface inflow from the drainage area downstream from the gaging stations was not measured. Much of this area is under cultivation and little runoff results from light rains. Estimates were made of unmeasured inflow from storm rainfalls of more than 0.5 inch.

OUTFLOW

There was no surface outflow from Lake Colorado City during the period of observation. The reservoir has not filled to spillway level since the dam was constructed in 1949, nor has any water been released since hat time. Provision was made to obtain a rating for he uncontrolled service spillway by relating discharge aeasurements, to be made below the dam, to the lake levation, in the event that spill occurred.

Periodic estimates of surface seepage below the dam, ndicate that this loss never exceeded 0.1 acre-foot per lay, and it was therefore disregarded. The possibility 'f deep seepage losses of any appreciable magnitude was onsidered remote. Hydrologic studies to be described ater confirmed that there were no measurable unac-:ounted-for losses from the reservoir.

CHANGE IN RESERVOIR CONTENT

The area and capacity curves used to determine shange in contents of Lake Colorado City were furnished <y Texas Electric Service Co., Fort Worth, Tex. (fig. 8).

?;QTJEE 8.—Area and capacity curves for Lake Colorado City, Tex. (Furnished by Texas Electric Service Co.).

Ihe curves are based on areas computed from four aerial p" otographs of the reservoir area, taken with lake sur- *•f,c&* at elevations of 2,047.3,2,053.3, 2,061.2, and 2,068.2 :cet above mean sea level. The aerial photographs were yirrected for length on the basis of transmission alinenent maps. Available also was the area at the 2,070- Vot contour, determined from deed records based on ield surveys.

The area and capacity curves used in the computa-Ions agree closely with previous curves, furnished by the engineering firm that designed the dam, ai d are accepted as accurate. The surface area of the canal ∇ as considered insignificant and was not taken into account.

Changes in lake stage were measured with a Stevens A-35 waterstage recorder located in the Colorado City pumping station on the north shore of the lake about 1.5 miles north of the dam (fig. 3). The stilling well for the recorder was a 2-foot concrete pipe attached to the outside of the pumping station.

The staff gage attached to the stilling well was read to hundredths of a foot and the midnight elevations used in computing changes in stage were taken from the recorder chart. Figure 9 shows the monthly contents of Lake Colorado City.

WITHDRAWALS FOR MUNICIPAL USE

Withdrawals by Colorado City for municipal use were the only appreciable diversion from Lake Colorado City. Records of daily withdrawals were computed from pumping records furnished by Colorado City Water Department. Figure 10 shows the total monthly diversion by Colorado City. Withdrawals averaged 2.4 acre-feet per day and ranged between 0 and 5.8 acre-feet per day. As the total of 1,130 acre-feet withdrawn amounted to less than 9 percent of the computed evaporation over the period of observation, even a 10 percent error in the pumping record, which is not likely, would not have an appreciable effect on the results. Consumptive use by the powerplant and by residents along the lake shore was minor in comparison to the diversion to Colorado City and was neglected.

RAINFALL ON THE LAKE SURFACE

Rainfall was measured in four 8-inch nonrecording gages located around the lake (see fig. 3). A simple average of the catches of the four gages was taken as the mean rainfall on the lake surface. A Weather Bureau recording gage located at the powerplant was used to determine the time and duration of the rainfall.

Figure 11 shows the monthly rainfall for the four nonrecording gages. Total average rainfall on the lake surface during the period of observation was 14.8 inches, which added 1,562 acre-feet to the reservoir. Rainfall of 0.05 inch or more occurred on 41 days during the 468 days of observation. On 10 days the rainfall was more than 0.5 inch; on only 1 day was there more than 1 inch of rain.

Although the total rainfall on the lake surface amounted to 13 percent of the measured inflow and was a significant item in the water budget, there were many periods in which no rainfall occurred.

ENERGY BUDGET

Net solar and atmospheric radiation at Lake Colorado City were measured by the Cummings radiation integrator (CRI), a pan of water heavily insulated to minimize heat transfer through the sides and bottom. Records obtained during the Lake Hefner studies (Harbeck, 1954) and the Lake Mead studies (Koberg, 1958) showed that the CRI gave accurate results and was a reliable substitute for the conventional radiation

FIGURE 11.—Monthly rainfall at Lake Colorado City, Tex.

GEOLOGICAL SUKVEY PEOFESSIONAL **PAPER** 272 **PLATE 1**

A. INSTRUMENTS LOCATED NEAR THE TEXAS ELECTRIC SERVICE CO. PLANT

B EARLY MODEL OF THE WHITNEY UNDERWATER THERMOMETER

SMALL FLOAT USED TO SUPPORT SIX'S MAXIMUM-MINIMUM THERMOMETER

f quipment, which is expensive and requires much more *i* ttention.

RADIATION MEASUREMENTS

The CBI was located near the Texas Electric Service Co. intake structure where 110-volt alternating current vas available. Daily inspections and occasional servicing were made by plant personnel. Figure 3 shows the location of the CRI installation. A complete c escription of the CRI is presented in the Lake Hefner studies by Harbeck (1954, p. 120). The CRI was placed in operation on July 21, 1954, and was cont'nued in operation until October 31, 1955. Plate 1A ϵ ows the completed CRI installation.

The CRI was serviced at approximately 10-day intervals, using the technique developed at Lake Hefner (Harbeck, 1954, p. 121-122) except that plant personnel occasionally added water, when the water hvel in the pan was below an arbitrarily selected reference point, which was about 1 inch below the reference point gage. The volume and temperature c f the water added were measured.

The recording of temperature data was the principal nainterance problem, and it is discussed in the section en mass-transfer instrumentation.

TEMPERATURE PROFILES OF THE LAKE

Temperature profiles were obtained with a Whitney mderwater thermometer, which uses an electrical rircuit with a small battery to supply the power recuirements. The circuit is a Wheatstone bridge, of rhieh the underwater sensing element is one arm. 'nhe underwater sensing element is a small thermistor rhich responds rapidly to changes in temperature.

Temperatures are read directly from a calibrated 3?ale. The depth of water is measured by the length : f line from the sensing element to the water surface. Γ ate 1*B* shows one of the earlier models of the Whitney mderwater thermometer.

Although the instrument is nonrecording, it is portable, and for some purposes is therefore preferable to the bathythermograph (Spilhaus, 1938) and the temperature profile recorder (TPR) (Anderson and Purke, 1951) because it is portable. The bathythermograph winch and the TPR are not portable and require r semi-permanen t installation in the boat.

The calibration of the Whitney underwater thermometer was checked just before the investigation regan and immediately after completion. The cali ration checks agreed within 0.1°C, which is satis-^ctory for evaporation computations. Occasionally ield checks were made by placing the sensing element "* a bucket of water and measuring the temperature f the water with a calibrated thermometer. These :hecks always agreed within 0.1 °C.

The only maintenance problem of the Whitney underwater thermometer occurred when the sensing element was broken during a sounding and had to be returned to the manufacturer for repairs.

MASS **TRANSFER**

The mass-transfer instrumentation at Lake Colorado City was considerably simpler than that used at L^ke Hefner and Lake Mead. Equipment was provided for measuring wind speed and wet- and dry-bulb temperatures at only one height, and for measuring weter surface temperatures. Copper-constantan thermocouples were used for most temperature measurements.

The thermocouple voltages were recorded on a Minneapolis-Honeywell multiple-record Brown recording potentiometer. The recorder used has a strip type chart and a print wheel for printing six different records in sequence at 30-second intervals. The calibrated accuracy of the chart scale is within ± 0.02 mv (millivolt) or $\pm 0.5^{\circ}$ C and is sensitive to a change of 0.004 mv or 0.1°C. Power is supplied from a 110-volt alternating current source. The thermocouple psychrometer (Bellaire and Anderson, 1951) used successfully at Lake Hefner and Lake Mead was also used at Lake Colorado City.

Wind speed was measured with a standard 3-cup Robinson-type contact anemometer. A 4-digit electric reset counter was used to record wind movement. The counter was read daily and the dial reset to zero.

Two Six's maximum-minimum thermometers were used to study the areal variation in water-surface temperature. These thermometers were supported by floats just below the water surface with a small radiation shield above the thermometer bulb, as shown in Plate 2. Observations of maximum and minimum water-surface temperature were made once each day.

The thermocouple psychrometer and the anemometer were mounted on a mast attached to the vertical member at the left of the ladder leading to the top of the plant intake structure, as shown in Plate 1A A float was anchored in the lake near the intake structure with a thermocouple mounted just below the water surface. Two thermocouples were mounted on the CRI, 1 for measuring the water surface temperature in the CRI and 1 for measuring the temperature of the overhanging rim. The thermocouples were all connected to the Minneapolis-Honeywell recorder located near the CRI. Figure 3 shows the location of this installation which was placed in operation on July 20, 1954. and was continued in operation until October 31, 1955.

Figure 3 also shows the location of the two floats with the Six's maximum-minimum thermometers. One is located near the dam and the other in midlake. It would have been desirable to have had one of the floats nearer the north end of the lake, but this part of the lake is open to the public and vandalism was feared. These floats were installed on July 20, 1954, and temperatures will be observed until the need for these data no longer exists.

In addition to the mass-transfer instrumentation, a standard Weather Bureau class A evaporation pan was installed near the CRI by Weather Bureau employees. Figure 3 shows the location of the class A pan which was placed in operation on July 15, 1954.

PERFORMANCE AND MAINTENANCE OF EQUIPMENT

ACCURACY CHECKS

Periodic checks of the equipment in the field were made to maintain prescribed accuracy requirements. The temperature and humidity data were checked 2 or 3 times a week with a calibrated sling psychrometer. The sling psychrometer readings agreed with the thermocouple psychrometer records within 0.5°C. After dust storms, greater discrepancies were noticed, but after the wet-bulb wick was cleaned the readings agreed within 0.5°C.

The anemometer assembly was replaced and overhauled at intervals of about 30 days.

The thermocouple reference bath was contained in a 4,300 milliliter (ml) vacuum flask. It was essential that a temperature of 0°C be maintained at all times. During the first summer, the reference temperature was found to be above 0°C on several occasions. After this was noted, the reference bath was inspected daily and ice added when necessary.

The Six's maximum-minimum thermometers were checked during thermal surveys by holding a calibrated thermometer just under the water surface and comparing with the indicated temperature of the Six's thermometer. These checks agreed within 0.3°C.

USABLE DATA

Table 1 shows that the yield of usable data was excellent. The maintenance crew was inexperienced in maintaining and servicing this type of equipment, but quickly learned to remedy any malfunction and is to be complimented on the excellent yield of usable data.

MAINTENANCE PROBLEMS

The Minneapolis-Honeywell recorder, which was the heart of the temperature-recording scheme, performed exceptionally well. No loss of data could be attributed to any malfunction of the recorder. The recorder required very little servicing and maintenance compared with the recording system used at Lake Hefner and Lake Mead, but of course it requires 110 volt alternating

current power, which was unavailable at the other places.

TABLE 1.—*Percentage of usable data obtained from meteorologica equipment*

Month	Temper- ature	Humidi'y	Wind	Maximun minimum water sur- face tem- perature
1954 September October November December	100 75 80 84 91 100	82 170 65 1 ¹ 93 74	91 90 77 90 93 94	
1955 January February April May August__________________________ September October	100 100 100 100 100 100 100 100 100 100	120 75 94 87 1Μ 87 100 100 100 100	81 100 94 100 90 90 81 94 97 97	10 1t

The wet-bulb reservoir was filled twice a week an was found dry twice during the period of operatior In the winter on cold days, the wet-bulb reservoi was frozen for a few hours. The main maintenanc problem in connection with the wet-bulb was keepin the wick clean whenever dust was blowing. Durin these periods, the maintenance crew inspected an cleaned the wick daily or twice daily if needed.

The anemometer operated satisfactorily and the los of wind data was attributed to having an inadequat battery for the electrical counter.

The Six's maximum-minimum thermometer require very little maintenance except for removing algae The loss of maximum-minimum temperature data was caused by inability to read and reset the thermomet< indexes each day, as sometimes occurred when the bot was temporarily out of service or when the lake $w\epsilon$ unsafe for navigation because of high winds.

The lake-surface thermocouple assembly require very little maintenance. During the first 4 months operation, trouble resulted from the insulation on tl thermocouple leads cracking and causing the leads 1 short, but it was remedied by installing thermocoup wire with better insulation.

SUMMARY

The instrumentation program at Lake Colorad City was considerably simpler than that used at Lal Hefner and Lake Mead with no loss of accurac The program effected savings in manpower and equiment at a considerable savings in cost.

The Whitney underwater thermometer performe exceptionally well. Thermal surveys were ma< quickly and easily.

The Minneapolis-Honeywell recording unit eliminated much of the maintenance problem encountered with the battery powered thermocouple amplifier used at Lake Hefner and Lake Mead. The initial cost was slightly higher, but the saving in manpower and loss of record was adequate compensation.

ENERGY-BUDGET STUDIES

The energy budget per unit area of a reservoir per unit time may be expressed as follows :

$$
Q_s - Q_r + Q_a - Q_{ar} - Q_{bs} + Q_v - Q_e - Q_h - Q_w = Q_s \qquad (1)
$$

in which $Q_s =$ solar radiation incident to the water surface

- *Qr=* reflected solar radiation
- *Qa=* incoming long-wave radiation from the atmosphere
- *Qar=* reflected long-wave radiation
- *Qbg=* long-wave radiation emitted by the body of water
- Q_i =net energy advected into the body of water other than that contained in evaporated water
- *Qe=* energy utilized by evaporation
- *Qh=* energy conducted from the body of water as sensible heat
- *Qw=* energy advected by the evaporated water *Q&=* increase in energy stored in the body of water

Conduction of energy through the bottom, heating due to chemical and biological processes, and transformation of kinetic energy into thermal energy are neglected because of their small magnitude. For a thorough discussion of each term in equation 1, the reader is referred to the report by E. K. Anderson $(1954, p. 74-110).$

For computational purposes, use is made of the following relations:

$$
Q_e = \rho_e E L
$$
; $Q_h = R$ Q_e ; and $Q_w = \rho_e c E (T_e - T_b)$

in which *E=* volume of evaporated water

pe— density of evaporated water $L =$ latent heat of vaporization *E=* the Bo wen ratio c =specific heat of water

Te= temperature of evaporated water

 $T_b=$ arbitrary base temperature

Substituting the above in equation 1 results in the following :

$$
E = \frac{Q_s - Q_r + Q_a - Q_{ox} - Q_{bs} - Q_\theta + Q_s}{\rho e [L(1+R) + c(T_e - T_b)]} \tag{2}
$$

The value of T_b , the base temperature, is immaterial provided that same base temperature is used in computing Q_{θ} and Q_{θ} , and provided further that a balanced water budget is used in making the computations.

The method of determining each of the quantities in equation 1 is described in the sections that follow.

NET INCOMING RADIATION

At Lake Hefner and Lake Mead solar radiation (Q_s) and long-wave radiation from the atmosphere (Q_a) were measured directly using conventional radiation instruments. Reflected solar radiation (Q_r) was measured directly at Lake Hefner using radiation instruments; at Lake Mead it was evaluated from relationship? developed during the Lake Hefner studies by Anderson (1954, p. 78). The reflected long-wave radiation was computed on the basis of laboratory determinations of emissivity and absorbtivity. For the purpose of determining evaporation it is not necessary to measure each of these items separately. Only their sum, which is the net incoming radiation $(Q_s - Q_r + Q_a - Q_a)$, is needed and that can be obtained conveniently using a Cummings radiation integrator (OKI).

Comparisons between the sum $(Q_s - Q_r + Q_a - Q_{ar})$ determined using the CRI and the sum of the same items measured separately, have been made using data obtained in more than 2 years of record at Lake Hefner and Lake Mead. The correlations were excellent, the maximum deviation being approximately 5 percent. Because of the limitations of the flat-plate radiometer, it is not known whether these deviations are the fault of the OKI or the flat-plate radiometer.

In this investigation the net incoming radiation was computed for each period of approximately 10 days between thermal surveys of the lake. The OKI was serviced each time a thermal survey was made. The computational methods used are almost identical to those used in the energy budget for the lake and which are explained subsequently.

The basic assumption is that the sum $(Q_s - Q_r + Q_a -$ *Qar)* is the same for a pan on the shore of a lake as it is for the lake itself. For short periods of time, say an hour or less, this assumption is probably not acceptable because of possible transient cloud effects, but for periods of a week or more it appears reasonable. Evaporation from the OKI is measured, and from this and other data, it is possible to solve equation 1 for the sum $(Q_s - Q_r + Q_a - Q_a)$. This computed value is then used in equation 2 to solve for evaporation from the lake.

RADIATION FROM THE LAKE

Long-wave radiation (Q_{bs}) emitted by the lake was computed using the Stefan-Boltzman law for blackbody radiation, with an emissivity of 0.970 for water as determined by Gier and Dunkle (Anderson 1954, p. 96-98).

The variation with temperature of long-wave radiation emitted by the lake is nearly linear over the range in water-surface temperature experienced in any energybudget period. For this reason an average value of long-wave radiation emitted from the lake for any period was computed from the average lake-surface temperature per period.

The average lake-surface temperature for each period was computed from the average daily lake-surface temperatures. The average daily lake-surface temperature is an average of the daily lake-surface temperatures at each float station. As the three float stations were located in the lower half of the lake, which was usually warmer than the upper half, a correction was applied to the daily average. This correction is based on the 26 lake-surface observations taken during the thermal surveys, and observations of lake-surface temperatures taken at each float station during the thermal survey. The correction varied from -0.1 ^oC in summer to -0.7 °C in winter.

BOWEN RATIO

The Bowen ratio, which has been widely used as a measure of the ratio of the energy conducted to or from the lake as sensible heat to the energy utilized for evaporation, is expressed as follows:

$$
R = \frac{\gamma (T_o - T_a) P}{(e_o - e_a) 1,000} \tag{3}
$$

in which $\gamma = a$ coefficient

 T_o =water surface temperature in $^{\circ}C$

Ta= air temperature in °C

P— atmospheric pressure in millibars

- e_o =saturation vapor pressure corresponding to the water surface temperature in millibars
- e_a =vapor pressure of the air in millibars

According to Bowen (1926) the value of the coefficient, γ , in the above equation varies between 0.58 and 0.66 but has a most probable value of 0.61.

The Lake Mead studies (Koberg, 1958) showed that the height at which both air temperature and humidity are measured makes very little difference insofar as the Bowen ratio is concerned. In the computations of Bowen ratios for Lake Colorado City the air temperature and humidity measured at the plant intake structure were used in all the computations. T_o and e_o were taken from the corrected average daily lake-surface temperature.

ADVECTED ENERGY

Advected energy (Q_v) is defined as the net energy gained by a body of water resulting from volumes of water entering and leaving the lake. It includes sub-

temperature of 0°C.

surface inflow and outflow and rainfall on the lake surface, and in the Lake Colorado City investigations it included the heat added to the lake by the powerplant. It does not include the energy contained in the evaporated water.

The various methods of determining inflow, outflow, change in storage, and rainfall are described in the instrumentation section. The flow diverted through the plant was taken from the Texas Electric Service Co. pump ratings. Occasional discharge measurements were made in the canal through which the water flows back into the lake (fig. 12). Based on these measurements, corrections were applied to the pump ratings, the maximum correction being a reduction of 7 percent in the pump rating. Actually, if no corrections had been applied the error in computed evaporation would have been less than 1 percent. The temperatures of these various items were computed in the following manner: temperatures of the inflow and the diversion to Colorado City were measured using a Stevens thermograph attachment to a water-stage recorder. Rainfall temperatures were assumed to equal the wetbulb temperature at the time the rain was falling, on the basis of data obtained at Lake Hefner (Harbeck, 1954, p. 123). Temperature of the water diverted to the plant for cooling and the temperature of the return flow to the lake were taken from plant records. Periodic checks were made of these temperatures using the Whitney underwater thermometer. These checks agreed within 0.5°C.

Advected energy was computed on a daily lasis. Density and specific heat were assumed constant for all computations. Figure 13 shows the advected energy for each month of the investigation.

ENERGY STORAGE

Energy storage (Q_0) in Lake Colorado City was computed from temperature profiles of the lake taken approximately every 10 days at 26 stations located throughout the lake. The selection of these 26 stations was based on a survey made May 4,1954, at 45 stations.

For each thermal survey the lake was divided into layers 3 feet thick. The temperatures in each layer were averaged to obtain the mean for the layer. The energy content in each layer was computed using: the mean temperature and the volume. The sum of these values gave the total energy content above an arbitrary base temperature of 0°C. Density and specific heat were considered constant, as in the computation of

advected energy, Figure 14 shows the energy storage by months.

SUMMARY OF MEASUREMENTS OF ENERGY-BUDGET TERMS

Table 2 was prepared to illustrate the relative magnitudes of the various items in the energy budget for Lake Colorado City and to show their seasonal variation. The sum $(Q_s - Q_r + Q_a - Q_a)$ was computed from the CRI data. Q_{bs} , the long-wave radiation emitted by the water surface, is an important term in the energy budget, and it is obvious that an accurate determination of average water-surface temperature for the entire lake is necessary. Except during infrequent periods of

storm inflow, the magnitude of Q_{\bullet} , the advected energy term is determined primarily by the amount of heat added by the powerplant. During the period July 21, 1954, to July 19, 1955, the amount of heat added by the powerplant averaged 59 cal cm^{-2} day⁻¹.

RESULT OF STUDIES

The figures of evaporation from Lake Colorado City using the energy-budget method are given in table 3. For the 363-day period, July 21, 1954, to July 19, 1955, evaporation was 92.78 inches, or 10,057 acre-feet equivalent to 93.3 inches, or 10,100 acre-feet, for a full year.

TABLE 2.—Average values, by periods, for each term in the energy budget for Lake Colorado City, in calories per square centimeter *per day*

Period		$Q_{a}-Q_{r}+Q_{a}-Q_{ar}$ Qы		Q ₀	Q.	Qı	Q.,	Qg
From-	$_{\rm To-}$							
July 21, 1954 ₋₋₋₋₋₋ $\rm{July}~30$ ---------- Aug. 9 Aug. 19 ------------ Sept. 1 Sept. 13 Sept. 23 4.1.1.1.1.1.1.1.1 Oct. Oct. $15 - 22 - 22 - 22$ Oct. $26 - 7 - 7 - 7 - 7 - 7$ Nov. 5 ------------ Nov. 16 ----------- Nov. 26 ₋₋₋₋₋₋₋₋₋₋₋₋ Dec. 6 ------------ Dec. 21. ---------- Jan. 3, 1955 Jan. 13 $Jan. 24$ ----------- Feb. $\text{Feb. } 14$ ----------- Mar. 1_{max} Mar. 11------------ Mar. 23 Apr. 4 ------------- Apr. 15 ------- Apr. 25 ---------- $\text{May } 5$ ------------ $\text{May } 16$ ---------- $\text{May } 26$ ----------- June 6 ------------ June 17 June 27. July 7 $\text{July } 19$ $July 29$ ----------- Aug. 19 Aug. 29. Sept. 8 ------------- Sept. 19 ₋₋₋₋₋₋₋₋₋₋₋₋ Sept. 29 ₋₋₋₋₋₋₋₋₋₋₋₋	July 30, 1954 Aug. $19 - 19 - 19$ Sept. 1 Sept. 13 Sept. 23 Oct. 4 Oct. 15 ------------ Oct. 26 ------------ Nov. 5 ------------ Nov. 16 ----------- Nov. 26 ₋₋₋₋₋₋₋₋₋₋₋ Dec. 6. Dec. 21 Jan. 3, 1955 Jan. 13 Jan. 24 $\text{Feb. } 4$ $\text{Feb. } 14.$ Mar. $1 - 1 - 1 - 1 - 1 - 1 - 1$ Mar. 11 ----------- Mar. 23. Apr. 4 Apr. 15. Apr. 25 ----------- $May 5$ ----------- May 16. May 26 June 6 June 17 ₋₋₋₋₋₋₋₋₋₋ June 27 ₋₋₋₋₋₋₋₋₋₋ July 7------------ July 19 $\text{July } 29$ ------------ Aug. $19 - 19 - 19$ Sept. 8 ------------ Sept. 19 Sept. 29 Oct. 20 ------------	1,448 1,512 1, 478 1,336 1,330 1, 267 1,213 1, 179 1,087 948 881 924 917 843 750 736 813 831 876 931 1,084 1,083 1, 142 1,182 1, 241 1,328 1, 227 1,350 270 1,375 1, 476 1,501 1.420 1,386 1, 472 1.350 431 1,339 1,293 1, 213 1,112	944 940 937 930 928 912 886 883 848 804 784 775 761 739 724 719 708 708 726 726 760 777 757 779 814 843 878 873 873 885 907 914 931 929 936 937 929 912 906 901 879	82 66 74 74 72 58 57 57 54 57 48 46 50 51 49 53 51 53 51 50 55 51 55 65 68 74 177 167 53 68 76 66 83 118 79 56 92 53 57 57 137	617 650 674 494 546 511 388 474 365 384 98 254 259 248 75 85 163 82 208 166 303 356 429 402 436 475 322 526 405 554 568 677 550 481 535 512 540 554 454 365 353	$-28\,$ 17 -34 -17 -3 -10 -18 24 32 101 11 18 -7 25 $\bf{0}$ 27 38 -15 48 $\boldsymbol{2}$ -57 10 25 -34 -113 -106 6 20 -22 -10 -37 -52 7 7 -3 9 4 24 -- 6 7 32	31 32 33 24 26 24 16 20 13 11 3 6 6 5 1 2 2 1 4 3 7 9 9 10 13 17 13 21 16 23 25 31 26 23 26 25 26 25 20 16 14	-34 -61 -58 -21 -95 -112 -2 -165 -117 $-\sqrt{295}$ 33 -83 -52 -123 -1 -44 -47 108 -59 84 126 -18 -23 90 159 173 185 77 51 -9 89 -3 - 11 64 57 -77 24 -123 -24 -19 -29
0 ct. 20 - - - - - - - - - - - - July 21, 1954 ₋	Oct. 31 July 19, 1955	1,092 1, 142	830 827	41 66	453 379	41 -- 4	15 14	-206 -8

FIGUEE 14.—Monthly energy storage in Lake Colorado City using a base temperature of 0°C.

TABLE 3.—*Evaporation from Lake Colorado City as computed for energy-budget periods, July 21, 1954, to October 31, 1955*

Period	Number of days		Evaporation	
$From-$	T ₀	in period	Inches	Acre-feet
r_{1} ly 21, 1954	July 30, 1954	9	3.77	485
\overline{y} 30	Aug. 9.	10	4.41	551
Aug. 9	Aug. 19.	10	4.57	556
Aug. 19.	Sept. 1	13	4.34	515
Rept. 1	Sept. 13	12	4.45	515
Rept. 13	Sept. 23	10	3.46	392
Boot. 23	Oct. 4	11	2.89	321
Oct. 4	Oct. 15.	11	3.51	383
Oct. 15.	Oct. 26	11	2.70	291
Oct. 26	Nov. 5	10	2.59	277
[*] 0V. 5	Nov. 16	11	.73	76
	\overline{N} ov. 26.	10	1.69	179
$\sqrt{6}$ v. 26	Dec. 6	10	1.74	182
	Dec. 21	15	2.48	256
Dec. 21 _ _ _ _ _ _ _ _ _ _ _ _ _ _ _ _	Jan. 3. 1955	13	. 65	67
'an. 3, 1955	Jan. 13.	10	. 57	58
an. 13	Jan. 24 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	11	1.20	124
an. 24	Feb. 4	11	. 59	60
	Feb. 14	10	1.39	141
¹⁷ eb. 14.	Mar. 1	15	1.66	168
" far. 1	Mar. 11.	10	2.01	201
" Tar. 11	Mar. 23	12	2.86	284
" " (ar. 23.	Apr. 4	12	3. 44	338
\pr. 4. ___ .	Apr. 15	11	2.98	288
Apr. 15.	Apr. 25	10	2,93	280
Apr. 25.	$\text{May } 5$	10	3. 20	302
" "ay 5	$May 16$ -----------------	11	2.41	236
Aay16 $^{\circ}$ / ay 26	May 26.	10	3.60	377
l'une 6.	June 6__________________ June 17.	11 11	2.99	330 442
T une 17.	June 27.	10	4.06 3.84	419
$\sin 27$	July 7	10	4. 59	492
" · uly 7	July 19	12	4. 48	471
	July 29 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	10	3.23	340
	Aug. 8.	10	3.69	397
ug. 8. .	Aug. 19.	11	3.80	403
'ug. 19_________________	Aug. 29.	10	3.66	389
'ug. 29	Sept. 8	10	3.74	392
9pt. 8	Sept. 19.	11	3. 37	349
ept. 19. .	Sept. 29	10	2.48	255
ϕ _p t. 29	Oct. 20.	21	5.00	584
et. 20. <u> - - - - - - - - - - - - - - - -</u>	Oct. 31. -----------------	11	3. 35	393
uly 21, 1954	July 19, 1955	363	92. 73	10.057
"ly 21, 1954	Oct. 31, 1955	467	125. 10	13, 559

WATER-BUDGET STUDIES

The inflow to Lake Colorado City during the 363-day period July 21, 1954, to July 19, 1955, was 6,422 acrefeet. The only outflow was the diversion to Colorado City for municipal purposes, which totalled 891 acrefeet. Using observed stages on July 21, 1954, and July 19, 1955, and the area table, the decrease in storage was 4,632 acre-feet, using the prismatoid formula. From these figures, water-budget evaporation was $6,422-891+4,632$, or 10,163 acre-feet, compared vith 10,057 acre-feet determined using the energy-budget method. Such close agreement is doubtless partly coincidental, for an error of only a few percent in measuring some of the water-budget or energy-budget quantities would make a much larger difference than was here indicated. For example, during the entire 16-month period, the energy-budget evaporation was 13,559 acre-feet as compared with water-budget evaporation of 15,061 acre-feet, a difference of approximately 10 percent.

The computations of advected energy as originally made were based to a large extent upon records for the gaging station Morgan Creek near Westbrook, Tex. Figures of energy-budget and water-budget evaporation during periods of storm inflow were not in good agreement for the period September 29 to October 20, 1955. During this period, runoff into the lake was extremely heavy; approximately 40 percent of the total inflow into Lake Colorado City in the 16-month period occurred during these 21 days. The difference between energy-budget and water-budget evaporation for this one period prompted a study of all major storm periods, which indicated that a substantial part of the inflow measured at the Morgan Creek gaging station did not reach the reservoir. Inflow to the reservoir was computed from change in contents of the reservoir during the storm period, diversions to Colorado City, and estimated evaporation losses. These last were computed using the mass-transfer formula, to be described in a subsequent chapter. Although the daily figures thus computed are admittedly subject to error, an error of even 100 percent would not affect the general conclusion. The results of the computations are given in table 4.

Table 4 indicates that there was a substantial reduction in flow between the gaging stations and the reservoir. Morgan Creek accounted for approximately 90 percent of the total volume of inflow during the periods studied, and this gaging station record was therefore scrutinized carefully to detect any possible errors. The stage-discharge relation for the gaging station on Morgan Creek is illustrated in figure 15. The peak stage attained during the rise of October 2,

FIGURE 15.—Stage-discharge relation for Morgan Creek near Westbrook, Tex.

1955, was 16.25 feet, well below the stages at which discharge measurements were obtained during the rise of May 11, 1954, shortly before the gaging station was established. The fact that the stage-discharge relation is so well defined by current meter measurements at high stages indicates that the reduction in flow between the gaging station and the reservoir is real and not the result of inaccurate stream gaging.

TABLE 4.—*Comparison between computed and measured inflow to Lake Colorado City during major storm periods*

Period	Diversion to Colorado City $(aore-t)$	Evapora- tion (acre-ft)	Gain in storage (acreft)	Computed inflow ¹ (acre-ft)	Measured inflow ² (acre-ft)
1955 May 11-13. May 23-26. July 28-30. Aug. 21-24. Oct. 1-7.	ŋ 10 11 13 4	78 196 87 125 227	2,011 1,996 818 369 3,918	2.096 2, 202 916 507 4,149	2, 251 2,534 1,303 729 5,579

> Sum of three preceding columns. * Sum of flow of Morgan Creek near Westbrook, Tex., flow of Graze Creek near Westbrook, Tex., rainfall on the reservoir surface, and a very small estimated volume of unmeasured inflow.

Another analysis was made using obsolete area and capacity curves to determine whether the discrepancy might result from errors in the stage-area relation. The differences between measured and computed inflow

were slightly less than those illustrated in table 4, but not enough less to indicate that error? in the area and capacity curves could be responsible for the differences between measured and computed inflow.

The possibility was considered that bank storage in Lake Colorado City might account for the discrepancy. If this were true, the water thus stored would be released during periods of zero inflow, when lake levels were falling as a result of evaporation and diversion to Colorado City. However, during these periods the agreement between water-budget and energy-budget evaporation was usually excellent, indicating that substantial unaccounted-for volumes of water were not reaching the reservoir, as would result if there were a return of bank storage. Two other periods prior to the time of the present study were also investigated. During both of these periods inflow was high, but of short duration, and the lake stage rose 7 feet or more in 1 day during each period. It did not appear likely that the entire bank storage capacity, if any, could be utilized in 1 day; it was believed that water would continue to enter into bank storage for at least a few days after the rise. This would result in a greater fall in reservoir stage than could be reasonably accounted for by evaporation and diversion. Such was not the case, however; the recession appeared to be normal after each rise. Since the studies indicated that appreciable volumes of water were neither being released from bank storage during periods of falling reservoir stages nor were being taken into bank storage immediately after sharp rises, it was concluded that bank storage in Lake Colorado City is negligible.

A reconnaissance was made of the Morgan Creek channel between the gaging station and the reservoir to determine whether geologic and hydrologic conditions are such as to provide an explanation of the channel losses. Above the mouth of Graze Creek, the average depth of the Morgan Creek channel is approximately 20 feet and its width is approximately 75 feet. Below this point the channel is somewhat wider, reaching perhaps 150 feet in a few places, and the banks are lower, the average depth being approximately 15 feet or less.

No evidence of any highly permeable beds was observed. Many outcrops of the Dockum group of Triassic age occur in the area, including the stream channel. The Dockum group along the eastern edge of the Llano Estacado to the south, in Borden, Scurry, Howard, and Mitchell Counties as described byHoots (1926) "has a total thickness of 300 to 450 feet and consists largely of dark-red clay with interbedded layers of gray cross-bedded sandstone and coarse sandstone conglomerate. The sandstone is invariably micaceous * * * The lower part of the Dockum group near Colorado City, in Mitchell County, though predominantly red clay, contains numerous beds of massive gray cross-bedded sandstone."

The outcrops seen in the Morgan Creek channel were substantially as described by Hoots. At one place well above the channel, but still within the flood plain of Morgan Creek, some poorly cemented sandstone was exposed; it was about 15 feet thick and the outcrop was not of great areal extent.

The geologic reconnaissance did not provide an explanation of the loss of water in the Morgan Creek channel. The alluvium in the stream channel appears to be quite permeable, and it is probable that during flood periods some of the water, apparently lost, enters the alluvium, to be returned later to the channel where it evaporates. Many such pools were observed during the reconnaissance on April 16, 1956, even though there had been no rain since April 5 when 0.66 inch was measured at San Angelo and 0.50 inch at Abilene. It does not seem reasonable that the entire loss could be accounted for *in* this manner, however. The Dockum group dips to the west, and any water entering the formation must drain away from Lake Colorado City. The same conclusion was reached on the basis of hydrologic studies of reservoir levels.

Since it appears that during periods of storm runoff there is a substantial loss of water in Morgan Creek between the gaging station and Lake Colorado City, inflow to the reservoir during the brief storm periods was computed from records of change in contents of the reservoir. Allowances were made for diversion to Colorado City and for evaporation. The latter was computed, but it was so small in comparison with inflow that even a 100 percent error in estimating it would have caused no significant change in the computed volume of inflow, which was used only ir the computation of advected energy.

With storm period inflow computed as described, the agreement between energy-budget and water-budget evaporation is considered excellent on an annual basis. For the individual periods, the agreement is not nearly so good, but the differences tend to compensate over longer periods and show no correlation with season. For short periods of time, errors in water-budget evaporation may be caused by errors in measuring mean lake stage. Only one water-stage recorder was used, and substantial errors may result from the fact that the reservoir surface is not level at times of high winds or high inflow. For example, if the stage is in error by 0.02 foot, the error in computing change in reservoir storage is approximately 25 acre-feet, an amount significant on a daily or even a weekly basis but not for longer periods. Errors from this scurce are not cumulative.

The close agreement between energy-budget and water-budget figures is corroborative evidence of the validity of the energy-budget method for the determination of evaporation from reservoirs. Results previously obtained at Lake Hefner and at Lake Mead showed this to be true, and the Lake Colorado City study confirmed the previous findings.

MASS-TRANSFER STUDIES

Data obtained at Lake Colorado City for the determination of evaporation using the mass-transfer method are: wind speed, wet- and dry-bulb temperatures, and water-surface temperature. No effort was made to place the anemometer and wet- and dry-bulb thermocouples at any standard height (such as 8 meters) above the water surface. For convenience these instruments were placed atop the superstructure for the travelling crane over the intake canal, plate *1A}* approximately 40 feet above the water surface, the exact distance depending on reservoir stage. Moreover the instruments were on shore, not on a barge in midJake.

It was anticipated that an empirical coefficient, valid only for this particular installation at Lake Colorado City, could be obtained for use in a simplified form of mass-transfer equation, as follows:

 $E=N n u (e_0-e_0)$ (4) in which *E=* total evaporation in depth units for period of *n* days u =wind speed

 $N=$ an empirical coefficient

*e0=*saturation vapor pressure correponding to the temperature of the water surface *ea=*vapor pressure of the air

In mass-transfer theory some equations have been derived to give evaporation at a particular point. Others have been proposed for the purpose of computing evaporation from the entire lake surface, thus taking into account the downwind decrease in evaporation. Practically all of the mass-transfer equations are of the same general type as equation 4. The coefficient *N* may be replaced by a complicated mathematical expression involving gamma functions or Pearson's function, and the wind term may have an exponent different from unity, but the general form of equation is the same.

For Lake Hefner a quasi-empirical equation of the same type as equation 4 was found to give good results (Marciano and Harbeck, 1954, eq. 58, p. 65). The effect of atmospheric stability was found to be insignificant, at least for figures of daily evaporation. At Lake Mead the value of *N* derived from the Lake Hefner data was found to give good results on an annual basis. For shorter periods of time the effect of atmospheric stability was clearly evident. An empirical stability adjustment parameter, roughly proportional to the Richardson number (Marciano and Harbeck, 1954, p. 52), was therefore used. The possible effect of atmospheric stability at Lake Colorado City could not be estimated in advance. It was anticipated that other seasonal effects might be noticeable. For example, both wind speed and vapor pressure might show a seasonal variation if there were a marked seasonal trend in wind direction, so that at certain seasons offshore and onshore winds might prevail. This would not present a problem if the instruments were mounted on a raft in midlake, but might with a shore installation as at Lake Colorado City.

ANALYSIS OF THE DATA

The relation between energy-budget evaporation and the product *n* u (e_o-e_a) is shown in figure 16. The energy-budget evaporation is for periods between thermal surveys, which averaged about 10 days in length. The wind speed, *u,* is the average daily wind speed in miles per hour; *n* is the number of days. The vapor pressure difference (e_0-e_a) is based on the average daily values of *e0* and *ea* during the period. All

FIGURE 16.—Relation between energy-budget evaporation and the product $n u (e_{o} - e_{a})$ using data obtained at Lake Colorado City, Tex.

the data were obtained at Lake Colorado City. The correlation between energy-budget evaporation and the product *n* u ($e_o - e_a$) is fairly good. The value of *N* to be used in equation 4 was found to be C 00251, obtained by dividing energy-budget evaporation for the period July 21, 1954 to July 19, 1955, (363 days) by *2n u* (e_0-e_a) for the same periods. The resulting equation is $E=0.00251 \; n \; u \; (e_o-e_a)$ (5)

in which *E* is in inches for the period of *n* days and *u* and (e_0-e_a) are computed as described above.

The data were smoothed by grouping 10-day energybudget periods in threes. Using equation 5, evaporation was computed for each group of 3 periods and compared with the energy-budget evaporation for the same group of periods. The residuals, or differences between the two figures, exhibit a rather poorly defined seasonal variation. Energy-budget evaporation is less than mass-transfer evaporation for the period October 26, 1954. to February 4, 1955. The reverse is true for the periods July 21, 1954, to October 26, 1954, and June 17, 1955. to September 19, 1955. The differences cannot be accounted for by the effects of atmospheric stability. In winter, water temperature is generally higher than air temperature. Because the lapse rate is unstable, evaporation should be greater than that given by equation 5, which is theoretically correct only for neutral stability. Thus the differences cannot be attributed to the effects of stability. It was previously suggested that a seasonal variation in wind direction might cause a seasonal variation in vapor pressure. Examination of the records at the Weather Bureau stations at San Angelo and Abilene shows no such variation, although prevailing winds at those stations are consistently from different directions.

Figure 16 indicates that the relation between energybudget evaporation and the product *n* u (e_{θ} - e_{θ}) may be closely approximated by a straight line, but that the line does not necessarily go through the origin. A computed least-squares line does have a small negative intercept (only 0.10 inch) on the evaporation axis, but the intercept is so small that it was disregarded, and evaporation is considered to be directly proportional to the product *n* u (e_o-e_a). A possible explanation of the intercept will be given later.

The use of equation 5 for computing evaporation from Lake Colorado City on a continuing basis would require that all of the mass-transfer instruments be operated indefinitely. In order to determine whether satisfactory results might be obtained using data already being obtained, analyses were made substituting data normally collected in connection with routine plant operations and at nearby Weather Bureau stations.

In the first of these, the intake water temperature at the plant was used to determine e_o . The vapor pressure of the air, e_a , was obtained from Weather Bureau records at Big Spring, Tex. The wind speed, *u,* was taken from the anemometer record at Lake Colorado City. The resulting equation was

$$
E=0.00260 \; n \; u \; (e_o-e_a) \tag{6}
$$

Another analysis was made using both wind speed and vapor pressure of the air measured at the Weather Bureau station at San Angelo, Tex. Other data were the same as in equation 6. The new equation is

$$
E=0.00201 \; n \; u \; (e_0-e_a) \tag{7}
$$

The relation between energy-budget evaporation and the products *n* u ($e_{o}-e_{a}$) as computed from the data described above is shown in figures 17 and 18.

All three figures, 16, 17, and 18, indicated a negative intercept on the evaporation axis. This is partly owing to the computational methods used. In figure 16, the wind speed, *u,* is the average wind speed for the period. The saturation vapor pressure, e_o , is that corresponding to the average water surface temperature for the period. The vapor pressure, *ea,* is the average vapor pressure for the period. The product, therefore, is the product of average values. A similar least squares computation was made using the average of the daily values of each product instead of the product of the average values for a period of a year. The intercept was found to be -0.03 inch. Although the intercept was negligible using the least squares technique, the conclusion is misleading, as a result of minimizing the deviations in the vertical direction. The data used for the independent variable in figures 16, 17, and 18 are not the

FIGURE 17.-Relation between energy-budget evaporation and the product *n* u (e_o-e_o) using data obtained at Lake Colorado City and Big Spring, T^{*}.

FIGURE 18.—Relation between energy-budget evaporation and the product *n u* (e_0-e_4) using data obtained at Lake Colorado City and San Angelo, Tex

same, and a negative intercept is indicated in each case. It must be concluded that the negative intercept results from errors in the energy-budget evaporation. A negative intercept, if real, would indicate that substantial volumes of seepage were entering the lake, even during prolonged dry periods. Because of the excellent agreement between energy-budget and water-budget evaporation during dry periods, this hypothesis must be disregarded.

The comprehensive tests of the CRI made at Lake Mead (Koberg, 1958) indicate that although on the average the CRI measured net radiation received by the lake very accurately, there is a slight seasonal bias. Net radiation as measured by the CRI is slightly less than that measured by conventional radiation instruments in winter and slightly greater in summer. The maximum deviation was approximately 5 percent. Studies are underway to eliminate the seasonal bias, if possible. It is believed to result, at least in part, from the variation with sun angle in the effect of the overhanging rim. Although the error from this source is small, it could account for computed figures of energy budget evaporation being slightly smaller in winter and larger in summer than computed figures of mass-transfer evaporation.

The seasonal bias in net radiation measured by the CRI, is too poorly defined to permit any corrections on the basis of the Lake Mead studies. It was therefore concluded that the negative intercept indicated by figures 16, 17, and 18 is not real but probably results from small errors in measuring net radiation. The applicable value of the coefficient to be used in equations 5, 6, and 7 was therefore determined from annual mean values of energy-budget evaporation and the product *n u* $(e_{\mathfrak{a}}-e_{\mathfrak{a}})$.

RESULT OP COMPUTATIONS

Evaporation was computed using equations 5, 6, and 7 for energy-budget periods of approximately 10 days each. The results are shown in table 5; figures of energy-budget evaporation are also given for comparative purposes. The seasonal bias is evident, for energybudget evaporation is consistently less than masstransfer evaporation during winter, and vice versa in summer. Despite the bias, the average difference without regard to sign is only 11 percent for equation 5, 18 percent for equation 6, and 15 percent for equation 7. Because the difference is believed to result in part from a slight seasonal error in measuring net radiation, the error is believed to be less than that given above. The percentage errors given are for periods of approximately 10 days; the comparable error for a period of a month is approximately 8 percent for equation 5 and slightly more for the other 2 equations.

There appears to be little difference in the accuracy of the results obtained using the various mass-transfer equations. For figures of monthly evaporation, the average error without regard to sign (assuming the energy-budget figures to be correct) can be taken to be approximately 10 percent for all 3 equations. Using equation 6 or 7 would eliminate the need for any psychrometric observations at Lake Colorado City. Equation 6 requires wind measurements at the reservoir, but operation of a totalizing anemometer presents no problem. Although equations 6 and 7 do not provide 2 completely independent estimates because some of the same data are used in each, they do use different wind and humidity data, and an average of the results obtained with the 2 equations might we'l be considered the best estimate of evaporation obtainable with a minimum of instruments, observations, and computations.

TABLE 5.—*Comparison between energy-budget evaporation and mass-transfer evaporation computed from three empirical formulas*

	Period	Energy- budget evapora-	Mass-transfer evapora- (inches) tion			
$From--$	$_{\rm To-}$	tion (inches)	eq 5	eq 6	eq 7	
July 21, 1954.	July 30, 1954	3. 77	3.46	3.26	3.59	
July 30 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	Aug. 9.	4.41	4.52	3,88	4.06	
Aug. 9.	Aug. 19.	4.57	4.35	4. 13	4.49	
Aug. 19.	Sept. 1.	4.34	3.73	3. 17	4.24	
Sept. 1 ₋₋₋₋₋₋ -----------	Sept. 13. <i>. .</i>	4.45	4. 11	3. 87	4.10	
Sept. 13.	Sept. 23 ₋ ---------------	3.46	3. 64	3.84	3.54	
Sept. 23.	Oct. 4 ₋ ---------- <i>------</i> -	2,89	2.95	2.98	2.82	
Oct. 4	Oct. 15.	3. 51	3. 51	3.67	3.07	
Oct. 15 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	Oct. 26	2.70	2.47	2.52	2,67	
Oct. 26.	Nov. 5.	2.59	2.78	3. 16	2.57	
Nov. 5.	Nov. 16.	. 73	1.00	1.08	1.16	
Nov. 16.	Nov. 26.	1.69	1.65	1.89	2,39	
		1.74	2. 11	2.46	1.93	
Nov. 26.	Dec. 6.					
Dec. 6 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	Dec. 21	2.48	2.72	3.39	3.57	
Dec. 21 ₋₋₋₋₋₋ -----------	Jan. 3. 1955	. 65	1. 15	1.37	1.51	
Jan. 3, 1955.	Jan. 13.	. 57	71	. 83	. 81	
Jan. 13.	Jan. 24 ₋	1.20	1.26	1.87	1.38	
Jan. 24.	Feb. 4 .	. 59	1.20	1.45	1. 20	
Feb. 4	Feb. 14.	1.39	1.46	1. 71	1.62	
Feb. 14. .	Mar. 1.	1.66	1.72	2.18	1.65	
Mar. 1.	Mar. 11	2.01	1.79	1.78	1.66	
Mar. 11.	Mar. 23.	2.86	2.49	2.80	2.32	
Mar. 23.	Apr. 4	3.44	3.89	4.40	4.60	
Apr. 4.	Apr. 15.	2.98	2.98	2.72	2.94	
Apr. 15	Apr. 25	2.93	2.74	2.69	1.96	
Apr. 25	May 5.	3. 20	2.45	2, 28	2.72	
May 5.	May 16	2.41	2.91	2.65	2,53	
May 16 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋	May 26_______________	3.60	3.75	3.73	3.33	
May 26.	June 6________________	2.99	3.66	3.39	3.88	
June 6 ₋₋₋₋₋₋₋ -----------	June 17.	4.06	4.53	3.53	3.46	
June 17.	June 27.	3. 84	3.54	3.59	3.59	
June 27.	July 7	4.59	3.92	3. 61	4.07	
July 7	July 19 ₋ ----------------	4. 48	3. 44	2.87	3.19	
July 19.	July 29 ₋₋₋₋₋₋₋₋₋ --------	3.23	2.93	2.47	2.81	
July 29.	Aug. 8.	3.69	3. 15	2.63	2.94	
Aug. 8.	Aug. 19.	3. 80	3.61	3, 25	4.69	
Aug. 19.	Aug. 29.	3.66	3.21	2, 62	3. 15	
Aug. 29. <i>.</i>	Sept. 8.	3.74	3. 13	2.83	3.49	
Sept. 8 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	Sept. 19.	3.37	2.89	2,37	3.33	
Sept. 19 ₋₋₋₋₋₋₋₋₋₋₋₋₋₋₋	Sept. 29	2.48	2.09	1.78	2.21	
Sept. 29.	Oct. 20.	5.00	5.73	5.07	6, 17	
Oct. 20.	Oct. 31	3.35	3.38	3.02	2.72	
July 21, 1954 ₋₋₋₋₋₋₋₋₋₋	July 19, 1955	92.78	92.59	92.75	92.62	
July 21.	Oct. 31	125. 10	122, 71	118.79	124.13	

EFFECT ON EVAPORATION OF ADDING HEAT TO THE RESERVOIR

It is common practice to withdraw water from a reservoir or stream, use it for cooling in a steam powerplant, and return the heated water to the reservoir or stream. In the humid regions of the United States, the actual consumptive use of water is of little consequence. In the arid or semiarid regions of the West, however, the resulting increase in evaporation from the stream or reservoir must be considered. Moreover in these Western regions, the variability of streamflow is considerably greater than in the relatively well-watered East, so that natural streamflow cannot always be relied upon to provide adequate amounts of water during dry periods.

Water storage must therefore be provided, and frequently it has been found possible to locate a steam plant on the shore of an existing reservoir or natural lake. Occasionally it is necessary to construct a dam and reservoir solely for the purpose of providing a dependable source of cooling water.

The effect of adding heat to a lake or reservoir has been studied by many investigators, including Lima (1936), Throne (1951), and Harbeck (1953). Until the Lake Hefner studies (U. S. Geological Survey, 1954) continuous measurements of atmospheric radiation had not been made, and it was impossible previous to that time to evaluate properly all the terms in the energybudget equation, thus preventing a direct computation of the effect of the addition of heat.

Based on the Lake Hefner studies, Harbeck (1953) presented a general method for the determination of the effect on evaporation resulting from adding heat to a reservoir. The theory was derived by combining the energy-budget equation and a simplified mass-transfer equation.

BASIC THEORY

The energy budget for a reservoir is given in equation 1, page 15. For a reservoir to which heat is being added by a steam powerplant, it is convenient to consider that Q_{v} , the energy advected into the reservoir, is composed of three parts, namely: Q_c , the heat added by the powerplant, *Qet,* the energy added by volumes of water entering the lake as inflow, and $Q_{\nu o}$, the energy removed by volumes of water leaving the lake as outflow. Equation 1 then becomes

$$
Q_s - Q_r + Q_a - Q_{ar} - Q_{bs} - Q_s - Q_a - Q_w + Q_{si} - Q_{vo} + Q_c = Q_o
$$
\n(8)

The addition of heat to a lake will not affect the first four terms, $(Q_s - Q_r + Q_a - Q_a)$ as their sum is the net supply of energy received as solar and atmospheric radiation. Nor will it affect Q_{vi} . The amount of water actually consumed in the plant is usually negligible, but could easily be taken into account in the computation of *Qe,* which is the additional energy contained in the water returned to the lake over that contained in the water diverted from the lake. It is assumed that after equilibrium has been reached, over a long period of time, the increase in energy storage, Q_{δ} , is negligible. If the amount of heat added by the plant is suddenly increased, some of the additional heat would be utilized in increasing the amount of energy stored in the lake, and its full effect on evaporation would not be realized until equilibrium is reached. Conversely, if $484263 - 59 - 4$

the amount of heat being added were suddenly decreased, stored energy would be released, and evaporation would continue at an accelerated rate until ε new equilibrium temperature is reached.

Thus from equation 8 it follows that when heat is added to a lake

$$
(Q'_{\mathit{bs}}-Q_{\mathit{bs}})+(Q'_{\mathit{e}}-Q_{\mathit{e}})+(Q'_{\mathit{h}}-Q_{\mathit{h}})+
$$

$$
(Q'_{\mathit{w}}-Q_{\mathit{w}})+(Q'_{\mathit{so}}-Q_{\mathit{vo}})=Q_{\mathit{e}} \qquad (9)
$$

in which the unprimed symbols refer to the lake in its natural condition and the primed symbols to the lake after heat has been added.

At Lake Colorado City the term $(Q'_{\nu o} - Q_{\nu o})$ was negligible. The only outflow from the lake was that diverted to Colorado City for municipal purposes. The amount of energy removed was extremely small compared with other items in the energy budget, and the difference, $(Q'_{\nu o} - Q_{\nu o})$ resulting from the fact that the temperature of water diverted to Colorado City was increased slightly by the addition of heat from the plant was of no consequence.

For computational purposes the equation may be rewritten in a more simple form as

$$
\Delta Q_{bs} + \Delta Q_s + \Delta Q_h + \Delta Q_w = Q_c \tag{10}
$$

in which Δ indicates the increment from the unprimed to the primed figures.

COMPUTATION METHODS

The various terms in equation 10 can be computed as follows (units: calories per square centimeter per day)

$$
\Delta Q_{ba} = 0.970\sigma [(T_o' + 273)^4 - (T_o + 273)^4]
$$
 (11)

in which $\sigma = \text{Stefan-Boltzman constant}$ for black-body radiation $[=1.171 \times 10^{-7}$ calories per square centineter per (degree) * per dayj.

$$
\Delta Q_e = \rho E' L' - \rho E L \tag{12}
$$

in which ρ =average density of evaporated water ($\div 1$) gram per cubic centimeter)

- *E=* average daily evaporation in gramr per square centimeter per day (\div centimeters per day)
- and $L=$ latent heat of vaporization in calories per gram at $T_o (= 595.9 - 0.545 T_o,$ very closely).

$$
\Delta Q_b = R' Q'_{e} - R Q_{e} = R' \rho E' L' - R \rho E L \tag{13}
$$

in which R = the Bowen ratio $=\frac{Q_h}{Q_e} = \frac{0.61P(T_o - T_a)}{1,000(e_o - e_a)}$ (14) and *P=* atmospheric pressure in millibars

$$
\Delta Q_v = c\rho (E'T_o' - ET_o) \tag{15}
$$

in which c =specific heat of water ($\div 1$ calorie per gram per degree).

But from mass-transfer theory, assuming no change in wind speed and that the possible effect of changes in atmospheric stability resulting from an increase in water temperatures is negligible

$$
\frac{E'}{E} = \frac{e_o' - e_a}{e_o - e_a} \tag{16}
$$

 e_o =saturation vapor pressure at T_o in mb

*ea=*vapor pressure of the air in mb, determined from *Ta* and *Tu.*

Equations 10 and 16 make it possible to compute natural evaporation from Lake Colorado City, or the evaporation that would have occurred if no heat had been added by the powerplant. They also can be used to compute the evaporation that would occur if the amount of heat added by the powerplant is substantially increased. The effect of varying reservoir water content can also be investigated.

Because the variation of the terms in equation 10 with water-surface temperature is not linear, a direct solution combining the two equations is not practicable. A successive approximation technique is employed, as follows:

Suppose that it is desired to compute what would have been the natural evaporation, *E*, from Lake Colorado City had no heat been added by the powerplant $(Q_c=0)$. The following data are available:

- *Qe,* the amount of heat added by the powerplant
- *E',* computed evaporation, using equation 8, with a known amount of heat *(Qe)* being added by the plant
- *T'0,* the observed water surface temperature
- *Ta,* the observed air temperature
- *eaj* the observed vapor pressure of the air
- *P,* the standard barometric pressure corresponding to the altitude of Lake Colorado City

It can be reasoned that without the addition of heat, Q_c , the water surface temperature (T_o) , would be lower than that observed (T'_{o}) . As a first approximation, an estimated value of T_o is used to determine e_o from saturation vapor pressure tables, and *E* is computed from equation 16. Using these values of E and T_o in equations 11, 12, 13, and 15, the various terms in equation 10 can be computed, and their sum should be equal to Q_c . If not, other values of T_c must be assumed and the computations repeated until a check is obtained.

RESULT OF STUDIES

The relative magnitudes of the values of ΔQ_e , ΔQ_h , ΔQ_{w} , and ΔQ_{w} indicate the efficiency of the various physical processes in disposing of heat. Computations were made as described above for the period July 21, 1954, to July 19, 1955. During this period the powerplant added 59 cal cm⁻² day⁻¹ (equivalent to 1.3 billion kilowatthours per year of her.t added to the lake). For this period $\Delta Q_e = 34$, $\Delta Q_e = 15$, $\Delta Q_w = 2$, and $\Delta Q_{ba}=8$ (all in cal cm⁻² day⁻¹). Of the total energy added to the lake by the powerplant, 58 percent was utilized to increase evaporation, 25 percent was conducted to the air above the reservoir, 3 percent was carried away by the evaporated water, and 14 percent was radiated to the atmosphere.

The results of the energy-budget studies showed that for the same 363-day period, evaporation from the lake was 92.78 inches, equivalent to 93.3 inches for a full year. Expressed in units of volume, the annual loss was 10,100 acre-feet. Average content of the reservoir was 22,300 acre-feet, so that the evaporation loss was 45 percent of the water remaining in the reservoir after evaporation demands were satisfied. The computations described on page 26 yielded the result that $E'/E=1.099$ or that the evaporation loss directly attributable to addition of heat by the powerplant *(E'-E)* was 910 acre-feet. Natural evaporation expressed in depth units, as is customary, was 85.0 inches, or slightly more than 7 feet.

Average water-surface temperature during the 363 day period studied was 18.8° C (65.8°F). The computations show that if no heat had been added by the powerplant, average water-surface temperature would have been 18.0°C (64.4°F). The addition of heat by the powerplant thus had little effec* on the watersurface temperature. The effect en the thermal structure of the lake will be discussed in the chapter that follows.

EFFECT OF INCREASING THE AMOUNT OF HEAT ADDED

The preceding analysis outlined the method used to determine the water-surface temperature and evaporation loss that would have been observed if no heat had been added by the powerplant. During the 363-day period selected for study, heat was added by the powerplant at a rate of approximately 1.3 billion kilowatthours (kwhr) per year. Similar computations were made for rates of 1.5, 2, and 3 billion kwhr per year to determine the resulting water-surface temperatures and the increase in evaporation from the lake. The reservoir content was variously assumed to be 25, 50, 75, and 100 percent of capacity. The greatest increase in water-surface temperature and the highest percentage

increase in evaporation obviously would occur when large quantities of heat are added to a nearly empty reservoir.

The results are illustrated in figure 19. It has previously been determined that the average water surface temperature would have been 18.0°C if no heat had been added to the reservoir. If the reservoir had been full during the entire year, the average water surface temperature would have been 18.5°C; if it had been 25 percent full, the temperature would have been 19.4°C. The observed temperature was 18.8°C. If 3 billion kwhr were added during a year, which is more than twice the energy added to the lake during 1954-55, the average water surface temperature would be 21.2°C

HEAT ADDED TO RESERVOIR, IN BILLIONS OF KILOWATTHOURS PER YEAR FIGURE 19.-Relation between amount of heat added to Lake Colorado City and water surface temperature for various reservoir contents. Computations are based on meteorological limnological data obtained 1954-55.

if the reservoir were 25 percent full and 19.2°C if full. The relation between the amount of heat added and the temperature rise, as shown in figure 19, is not exactly linear, although it may appear to be, over the limited range shown. The error in straight-line extrapolation to a value of 6 billion kwhr per year would be less than 0.1°C if the reservoir were full.

The preceding computations were based on the assumption that incoming radiation, air temperature and humidity, and water-surface temperature during 1954- 55 were reasonably representative of average conditions at Lake Colorado City.

Similar computations were made for 2 selected shorter periods, 1 in summer and 1 in whiter, to investigate seasonal effects on the expected temperature rises and increase in evaporation for various amounts of heat added to the reservoir. The summer period, June 27 to September 8, 1955, was 73 days in length, and the winter period, December 21, 1954, to March 1, 1955, was 70 days in length. During the summer period reservoir contents averaged 71 percent of capacity, and during the winter period 68 percent.

Figure 20 indicates that the temperature rise resulting from any given quantity of heat added by the powerplant is approximately twice as great in winter as in summer. The percent increase in evaporation is 3 times as great in winter as in summer, but natural summer evaporation was 4.5 times as much as in winter, so that the actual increase in evaporation, expressed in volumes of water, is greater in summer than in winter. The fact that the temperature rise is greater in winter than in summer is of little practical consequence because even the increased water temperatures in winter are much lower than summer water temperatures.

Figure 21 illustrates the effect on evaporation of adding heat to the reservoir. The increase in evaporation, expressed in water volumes, is practically independent of the content of the reservoir. The heat added by the plant is disposed of principally by evanoration, back radiation, and conduction, all of which are surface phenomena. With other variables held constant, the total amount of heat thus disposed of is therefore

 $\overline{}$ 2000 $\overline{}$ INCREASE IN RESERVOIR EVAPORATION RESULTING FROM ingrease in reservoir evaporation resulting from
Addition of Heat by PowerPlant, in Acre-Feet
8 5 ADDITION OF HEAT BY POWERPLANT, IN ACRE-FEET 1954 $\overline{}$ added $\overline{}$ Hea $\overline{}$ \mathcal{L} $\overline{\mathcal{A}}$ \sim \mathcal{L} $E = 710$ $\overline{}$ oration, in acre
by plant, in bil
nours per year $E_f =$ forced evapora
Q_c=heat added by
of kilowatthou \overline{A} \vert / \vert of kilowatthours per year λ \mathcal{L} $\overline{}$ ——
1
ESER

FIGURE 21.-Relation between the amount of heat added to Lake Colorado City and the resulting increase in reservoir evaporation.

proportional to the surface area of the reservoir. The increase in evaporation expressed as the ratio E'/E increases markedly as the reservoir content decreases (see fig. 19), but at the same time the reservoir area is decreasing, so that the product of the increase in

evaporation (expressed in depth units) times the surface area is practically constant at Lake Colorad City.

Average inflow to Lake Colorado City for the period 1948-55 was 12,540 acre-feet. If the natural evaporation figure of approximately 7 feet during 1954-55 is considered representative of average conditions, natural evaporation (E_n) from a full reservoir would be 14,210 acre-feet. Evaporation resulting from the addition of heat by the plant (E_t) in 1954-55 was 910 acre-feet (see fig. 21). The amount of water diverted for municipal use by Colorado City was nearly 900 acre-feet during this same period. The diversion plus total evaporation loss from a full reservoir under these conditions would be 16,020 acre-feet, which is greater than average inflow. Contents would then decrease until equilibrium was reached, or when $E_t + E_n + 900 =$ 12,540. Using 7 feet as the amount of natural evaporation, $E_n = 7A$, in which A is surface area in acres. From figure 21, E_f =710 Q_c , assuming that it is not affected by reservoir area and contents as explained above.

$$
7A + 710 Qc + 900 = 12,540
$$
 (17)

Equation 17 and the area and capacity curves for the reservoir were used to compute the volume of

water that Lake Colorado City would contain under various assumptions concerning the volume of average annual inflow and the amount of heat added to the reservoir. The results are shown in figure 22. If inflow over a period of years was exactly 12,540 acre-feet per year (the average for 1948-55), and if the heat added by the plant were 1.3 billion kwhr per year, reservoir contents would remain constant at about 25,000 acre-feet. If the heat added were increased to 3 billion kwhr per year, the reservoir stage would drop 2 feet. If inflow averaged only 5,000 acre-feet, which is considerably below average (but much greater than the figure of 990 acre-feet recorded in 1951), and if the amount of heat added were increased to 3 billion kwhr per year, reservoir contents would in time decrease to about 2,000 acre-feet. Figure 22 indicates that the volume of inflow is of much greater importance than the amount of heat added in determining the stage that the reservoir will reach.

Another analysis was made to determine what would happen during an extreme drought. It was assumed that inflow was zero, and that the reservoir was full at the beginning of the drought period. The amount of heat added by the plant was taken as 1.3 billion kwhr per year, and the diversion to Colorado City as 900 acre-feet per year. At the end of 4 years, the reservoir would be practically empty. At the end of 3 years, storage would be reduced to approximately 4,000 acrefeet, or 13 percent of capacity. At this reservoir level the water surface temperature would be 20.4°C. Although reservoir contents might decrease from 100 percent to 13 percent of capacity, the rise in water surface temperature would only be about 2°C.

CONCLUSIONS

The effect on evaporation caused by the addition of heat to Lake Colorado City was to increase evaporation from 85 inches to 93 inches during the 1-year period selected for study. In terms of volume, natural evaporation was 9,190 acre-feet, and the forced evaporation was 910 acre-feet. The increase in evaporation was approximately equal to the volume of water diverted by Colorado City for municipal use. The water-surface temperature was 0.8°C higher than it would have been if no heat had been added. The rise in water-surface temperature (relative to the temperature that would be observed if no heat were added to the lake) is almost directly proportional to the amount of heat added, for a given surface area. This relation is approximately correct for temperature rises of a few degrees under conditions of low humidity.

Some engineering handbooks give estimated quantities of heat that can be disposed of by cooling ponds in terms of British thermal units (Btu) per square foot of

pond area per hour per degree Fahrenheit difference between the water and air temperature. The temperature-difference theory thus expressed is incorrect; the temperature difference should be that between the water-surface temperature in the pond if no heat were added and the surface temperature resulting from the addition of heat. The fact that the air-water temperature-difference theory is erroneous may easily be demonstrated. The average annual water-surface temperature of lakes that do not freeze is usually somewhat lower than average annual air temperature because of the cooling effect of evaporation. Suppose that just enough heat is added to cause the water temperature to equal the air temperature. According to this theory no heat could be dissipated, which is not correct.

Apparently some of the empirical formulas and rulesof-thumb were based on some work by Ruggles (1912). His results were expressed in terms of the air-vater temperature difference, but his temperature differences were approximately 35°F owing to the large amount of heat added. This use of the air-water temperature difference introduced little error. Normally, however, the difference between natural water temperatures and air temperatures is but a few degrees, so that when the amount of heat added is relatively small, the resultant temperature rise also is small and the error may be considerable.

Computations were made using data obtained at Lake Hefner during the interagency investigations of 1950-51. To cause a water-surface temperature rise of 1°F required the addition of only 5.8 Btu per square foot per hour as compared with 6.7 for Lake Colorado City.

In a report by Harbeck (1953) data were given for a hypothetical reservoir located in a place whose climate was similar to that of southeastern Colorado. Several different assumed values of average annual relative humidity were used. The amount of heat required for a water-surface temperature rise of 1°F ranged from 4.6 Btu per square foot per hour for an assumed relative humidity of 30 percent to 8.7 Btu per square foot per hour for an assumed relative humidity of 70 percent. As humidity increases, a larger portion of the heat added is disposed of by evaporation and conduction.

It is impossible to give any average figure of the amount of heat required to cause a water surface temperature rise of 1 ° that would be generally applicable to all lakes and reservoirs. For any particular lake, the climatologic data required may usually be obtained from published Weather Bureau records. Hydraulic data for the lake, if not available, must be estimated. The computations required have been described at α and since they are neither difficult nor laborious, it appears preferable that an estimate of the effect of adding heat be made for each reservoir concerned instead of using average values.

Previous studies were made using the Lake Hefner data to compare the increased evaporation resulting from the addition of heat to a reservoir with the amount of water that would have been consumed by a cooling tower in disposing of the same quantity of heat (Harbeck, 1953). Those studies indicated that the increase in water evaporated from a reservoir would be less than half the water that would be consumed by a cooling tower. This finding is strictly applicable only to Lake Hefner, but the computational procedures described herein can be used to make a similar comparison for any reservoir. A discussion of the economics of the cooling tower versus the cooling pond method of disposing of excess heat is beyond the scope of this report, but it is obvious that the saving in water that might be effected by the use of a cooling pond is only one of many factors that must be considered.

At Lake Colorado City it was found that the volume of forced evaporation is directly proportional to the amount of heat added regardless of the contents of the reservoir. The amount of heat added by the powerplant during the study period was equivalent to 1.3 billion kwhr per year. If this were increased to 3 billion kwhr per year, the forced evaporation would be increased in the same proportion. If the reservoir were full, the water surface temperature would be increased only 1.2°C, but if contents were only 25 percent of capacity, the temperature rise would be 3.2°C.

Expressed in units commonly used in engineering handbooks, the amount of heat disposed of in Lake Colorado City ranged from winter to summer between approximately 4 and 8 Btu per square foot per hour per degree difference in water-surface temperature. The average for the year was 6.7 Btu per square foot per hour per degree temperature difference. The temperature difference (in degrees Fahrenheit) is the rise in water-surface temperature, not the air-water temperature difference. These figures are not applicable to other reservoirs. Changes in humidity of the air have a marked effect on the temperature rise to be expected, thus also affecting the amount of heat disposed of, if it is expressed in terms of the temperature rise.

EFFECT OF ADDED HEAT ON THE THERMAL STRUCTURE

CIRCULATION IN THE RESERVOIR

In the preceding chapter we analyzed the effect on evaporation of adding heat to the reservoir. We also determined the water surface temperatures to be expected under various assumptions of amount of

heat added and of reservoir levels. It was unnecessary to consider how adding heat to the reservoir would affect water temperatures to be expected at various depths and locations in the reservoir, because the processes through which heat is returned to the atmosphere are surface phenomena, and temperatures at depth have no effect.

Because of its midlake constriction (see fig. 3), Lake Colorado City is probably not as efficient a "mixing bowl" as it would be if it were circular in shape. The question arises, whether for practical purposes mixing might be confined to the lower basin, thereby reducing the effective size of the lake.

The point at which the canal discharges heated Water is nearly a mile from the point at which water is withdrawn from the lake. Although the amount of water withdrawn depends on the powerplant load, most of the time the powerplant withdraws approximately 275 cfs from the lake. The temperature rise is approximately 4.8°C in summer and 3.7°C in winter.

Figure 23 shows the variation of water temperature with depth in the upper and lower besins for selected soundings in January and July. Although a few temperature measurements were made at depths greater than 40 feet (the lowest limit in figure 24) in the old stream channel near the dam, 98 percent of the water in the reservoir is above this depth at full pool. In July, water temperatures in the upper basin were only a few tenths of a degree different from those in the lower basin, but in winter, a difference of about 2° was observed. During fall and vinter the lower basin was definitely warmer than the upper basin, but during spring and summer little difference was noted.

The two principal factors causing mixing in the lake are density differences and wind. Wind speeds during February, March, and April are usurlly greater than in the remainder of the year, but very low wind speeds are uncommon at any season, and doubtless much of the mixing is caused by the wind.

Figure 23 indicates that the temperature difference between the upper and lower basins is substantially greater in winter than in summer. There is also a measurable difference in density between the two basins in winter although the difference is very small a maximum of 0.02 percent at low temperatures. This is shown in figure 24 where, using data obtained from the thermal surveys, mean water temperatures were computed for the upper basin and the lower basin and then converted to equivalent densities. At high temperatures the average densities in the two basins are almost exactly the same. At low temperatures the density in the upper basin is slightly greater than in the lower basin.

FIGURE 23.—Comparison between average water temperatures in upper and lower basins of Lake Colorado City, Tex.

1 000 .999 ^o *So* \lt .998 density of
}rams per ci 997 nperature of upper basin, in degrees centigrade 996 995
.995 .995 .996 .997 .998 .999 1.000 AVERAGE DENSITY OF WATER IN LOWER BASIN, IN GRAMS PER CUBIC CENTIMETER FIGURE 24.-Comparison between average densities of water in upper and lower basins, Lake Colorado City, Tex.

Little is known of the density differences required to initiate and maintain density currents, except that the differences may be extremely small. An approximate computation was made based on the following assumptions: (1) The irregular shape of Lake Colorado City was replaced by a trapezoidal trough having an

equivalent volume and surface area. (2) Warm from the powerplant was assumed to flow from the lower end to the upper end of the lake on the surface and return on the bottom. (3) The two layers of up-lake and down-lake flow were made equal in crosssectional area at the longitudinal center of the trough.

Using commonly accepted formulas for flow in open channels, it was found that the flow caused by a cufference in density between water at 20°C and at 19°C was substantially greater than the flow diverted by the powerplant for cooling purposes. The temperature difference of 1°C is the difference between the tenperature of the top layer and the temperature of the bottom layer. Such a temperature difference is commonly observed, as is illustrated by the temperature profiles of figure 23. In summer the difference is much greater than 1°. In winter the difference may bo less than 1°, but if the circulation is impaired by weak density currents, the effect is of no practical importance, since the water withdrawn at that time is always much colder than in the summer.

The possibility that the reservoir might be effectively divided into two parts at extremely low stages was investigated. If the two parts were connected only by a very small channel, it is conceivable that the upper part of the reservoir would not be effective in disposing of heat. The original stream channel traverses the entire reservoir bottom, however, and if the watersurface elevation is below the bankful stage of the old stream bed in the upper basin, there would remain no flooded areas of substantial size in the upper basin. The sluggish water in the old stream channel would remain, but its volume would be insignificant compared with the volume of water remaining in the lower basin. At extremely low stages, therefore, it can be reasoned that practically all the water remaining in storage, small though it may be, would be utilized in disposing of heat, providing the reservoir level does not drop below the powerplant intake.

The results of the theoretical analysis and the study of observed density difference between the upper and lower basins indicate that heat from the powerplant is being effectively dispersed over the entire lake. Although temperatures at one end of the lake may be several degrees different from those at the other end, density differences are very small

WITHDRAWAL TEMPERATURES

Previous analyses have determined the water-surface temperatures that may be expected under various assumptions of amount of heat added by the plant, and have indicated that the entire lake is being utilized to dispose of the added heat. If the amount of heat added to the reservoir should be increased, information is needed concerning the expected temperature of the water to be withdrawn from the lake.

The intake canal is quite deep in order that water may be withdrawn even at low reservoir levels. Water flows through the short intake canal to the pumps, and the volume of flow is sufficiently great to insure that the flow is turbulent at practically all times. The variation of temperature with depth in the intake canal appears to correspond exactly to that observed in the reservoir, and the average temperature of the water after it has been pumped to the plant is the same as the mean temperature of the water flowing in the intake canal. Although the pump intakes are near the bottom of the intake canal, the water is being withdrawn from all levels.

A study was made of the relation between intake water temperature and the average temperature of the entire lake, as determined from the thermal surveys. A least-squares regression analysis indicates that there is no significant seasonal variation in the relations between intake water temperatures, water-surface temperatures, and average lake temperatures. A simple comparison of average temperatures is adequate. From data obtained at the times of the thermal surveys made at approximately 10-day intervals during the entire 16-month period of observation, the average intake water temperature was found to be 20.38°C, the average water surface temperature was 20.33°C , and the average lake temperature was 19.91 °C. The indicated difference of 0.05° C between intrike temperature and water-surface temperature is probably instrumental error. The reason the average lake temperature is 0.47°C lower than the intake temperature is that approximately 15 percent of the water in the reservoir during 1954-55 was below the level of the intake. This bottom water is colder than the water above the intake, so that the mean temperature of the water between the intake level and the surface is slightly higher than the average lake temperature.

CONCLUSIONS

The studies of the effect on the thermal structure of adding heat to the reservoir indicate that the entire lake is being effectively utilized in disposing of the heat added by the powerplant.

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 $\mathcal{L}(\mathcal{A})$ and $\mathcal{L}(\mathcal{A})$. The set of $\mathcal{L}(\mathcal{A})$ $\label{eq:2.1} \frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^{2} \left(\frac{1}{\sqrt{2}}\right)^{2} \left(\$ $\label{eq:2.1} \frac{1}{\sqrt{2}}\int_{\mathbb{R}^3}\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2.$

APPENDIX

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 $\mathcal{L}(\mathcal{A})$ and $\mathcal{L}(\mathcal{A})$

EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY 37

TABLE 6.—*Daily lake stage and contents of Lake Colorado City, July 21, 1954, to October SO, 1955*

[Stage and content at 12:00 p. m. Add 2,000 feet to lake stage to convert to reservoir and mean sea level datum]

TABLE 7.—*Daily withdrawals from Lake Colorado City for municipal purposes, July 21, 1954, to October 30, 1955, in acre-feet*

	1954					1955										
Date	July	Aug.	Sept.	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.	Oct.
1. 1 10 -------- --------- 11 12 , , , , , , , , , , , , , , , , , , 14. 15 ₋₋₋₋₋₋₋ --------- 16. 17 18 ------------------ $19 - 12 - 12$ 21. $22 - 22$ 23 24 $25 - 20 - 20$ 28 29 $30 - 1 - 1 - 1$	1. ------------------ 1. 5.0 5.0 4.9 4.5 4.9 5.8 5.5 4.9 3.6 5.0	2.6 3.9 4.1 4.6 4.4 4.6 4.7 3.0 3.1 4.5 4.3 4.4 4.0 4,5 3, 2 4.8 4.3 4.5 4.5 4.2 3.7 3.8 4.3 4.4 4.7 4.1 4.4 3.4 3.3 3.8	3.6 3.3 3.5 3.2 2.8 3.3 1.8 2.1 1.8 2.5 2.6 2.4 3.5 3.8 3.8 3,3 2.8 3.0 2.6 3.7 2.9 2.8 2.8 2.6 2.8 2, 2 3.0 3.4 2.3 2.3	\cdot 8 2,1 1.7 2.7 2,0 1.1 1.3 1.8 2.1 1.7 2.3 2.1 2.8 1.6 1.6 1.8 1.4 1.6 1.9 1.7 1.9 1.5 1.5 1.1 1.1 1.4 1.5 .6 1,6 1.2	1, 3 .9 . 9 1.1 1.2 1.2 .9 1.4 1.2 1.2 1.3 1, 2 2, 2 .9 2.0 1.7 1, 8 .9 1.6 1.6 1, 4 2.4 1,3 2.1 0 0 $\bf{0}$ θ 0 0	$\bf{0}$ $\bf{0}$ 0 0 2.0 2.5 2.4 2.5 2, 6 2.3 2.5 1.5 1.8 1.4 1.8 1.8 1.9 1.8 1.4 2.8 2,1 1.8 2.3 1.7 2.2 2, 2 1.7 1.8 2.1 1.5	1.5 1.8 1.9 2.0 1.5 1.4 1.7 1.7 1.3 1.3 1.8 1.9 1.8 1.4 1.5 1.5 1.7 1.8 1.7 1.5 1.5 1.4 1.4 1.7 1.9 1.7 1.6 1, 8 1.4 1.8	1.9 1, 7 1.4 1.6 1.6 1,3 .6 0 2.2 1.5 1.7 1.7 1,0 2,4 3.0 2.0 1.7 1.6 1.4 1.3 2,0 1.7 1.7 1.6 1.9 2,0 2,3 2.4	2.6 1.9 2.9 2.8 2.0 1.4 2.4 2.4 2.7 2.6 3.2 2.6 2.1 3.4 3.0 1.5 2, 5 1.6 1.6 1, 8 1.4 1.5 1.9 2.2 1.6 1.4 1.8 2,1 2.5 1.7	2.2 3.1 1.9 1.5 2.7 2, 2 2.8 2.9 2.0 1.9 4.1 2.4 3.2 3.3 3.5 3.7 3.1 4.0 3.6 3.0 3.5 3.6 1.7 3.3 4.4 3.2 4.6 4.1 3.5 3.9	2.1 3.5 4.8 4.6 2.0 1.9 3.3 2.6 2.0 1.6 2.0 1.9 2.6 2.3 2,3 3,3 2.1 2.5 1.6 1.7 2,1 2.6 1.7 2,0 3.1 2.7 2.9 2.5 1.6 3.0	1.3 1.9 2, 8 2.7 2, 2 2, 6 3.8 2.5 1.0 2,6 2.0 2.4 4.3 3.9 4.0 2.5 2,7 3, 3 2,9 2.9 3.0 3, 8 4.4 3.9 3.9 3.6 4, 2 3.4 3, 7 4.4	4.6 4.0 2.5 3.4 4.3 4.3 4.2 4.2 4.3 3.5 4.1 3.4 2,0 2.8 2,0 3.0 1.7 1.4 1,9 1.9 3.5 3.3 4.1 3.3 3.6 3, 7 3.2 2.8 4.1 4.5	4.7 4.4 4.8 3,7 4,3 4.4 5.1 3.0 5.3 5.2 3.0 4.1 4.3 3.5 3.9 4.4 4.4 3.7 2.8 3.5 2.1 2.8 4.1 4.4 3.0 3,6 4.6 3.4 4.7 3,9	3.4 2.6 3.2 1.6 2.6 3.5 3.3 3.7 3, 3 3.2 .8 1.0 Ω 1,3 1.8 1, 2 1.7 1.7 1.6 2,6 2.7 1.8 1.2 .8 0 .9 0 .6 1.5 1.5	1.7 0 \cdot 7 1.2 $\bf{0}$ \cdot 3 $\bf{0}$ $\ddot{.}9$.6 0 1,0 0 .6 0 . 4 1.4 \cdot ⁴ 1.5 1.0 1.1 1.6 \cdot 9 . 6 1.1 0
$31.$	3.9	3.4		0.8		2.5	2.0		2,4		2.3	$- - - - -$	4.3	2.4		

TABLE 8.—*Daily rainfall in inches at Lake Colorado City, Tex.*

 $\left[\mathrm{T}\!=\!\mathrm{trace}\right]$

			Nonrecording gages				Nonrecordir ^o gages						
Date	1	$\mathbf{2}$	3	4	Average	Date		$\mathbf{2}$ \sim	3	4	Average		
1954 July 31 8. Aug. $30.$ 31 ----------- Sept. . Oct. $2-$ 5. 6. 25. ----------- 27. ----------- Nov. 14. Dec. 11 28. 1955 Jan. 14. ----------- 17. Feb. 3. 18. . . . Mar. 18 ------------ 20. ------------ Apr. 22 May 5.	0 .03 0 0 т .18 0 .13 0 0 0 .76 .38 .14 .25 .05 .26 .19 .01 .02 . 77 .02 ጥ .20 .01 .06 .85 .87 . 44	0.02 .23 .06 .01 т .18 0 .17 0 $\bf{0}$ T .61 .51 .12 .76 .04 .23 .20 .01 ጥ .50 .02 .01 . 41 .04 .05 .87 .94 .49	0 .10 т 0 .15 .10 .05 0 .10 .50 .60 .05 .27 0 .22 $\bf{0}$ \mathbf{T} .46 0 .01 .52 T .05 .47 .08 .82	0 т 0 .05 т 0 0 .21 0 0 1.10 .55 ጥ .31 0 .35 0 $\bf{0}$ т .24 $\bf{0}$ ጥ .75 T .06 .33 ጥ .47	0.005 .090 .015 .002 .050 .115 .012 .075 .052 .025 .000 .742 .510 .078 .398 .022 .280 .152 .005 .005 .492 .010 .005 .470 .012 .055 .630 .472 .555	1955 May 10. ----------- 11_____________ 16. $17-$ ------------ 19. . $20 -$ ----------- 23. $25 - 1 - 1 - 1 - 1 - 1 - 1$ June <u>.</u> 16. 19. $28 - 1 - 1 - 1 - 1 - 1 - 1$ July 12 Aug. 4 Sept. 4 $11 - 1 - 1 - 1 - 1 - 1 - 1$ 23. Oct. 1. 3.	0.35 .89 .18 .17 .16 .08 .41 .28 .62 .04 ,12 .32 . 53 0 . 45 .51 .18 .18 .33 0 .18 .65 .14 .22 1,72 .06 0 .41 .07	0, 28 .85 .39 .09 .08 .02 .47 .30 .64 .03 .06 .44 0 . 19 .06 .40 .54 .17 .40 .28 .10 .34 .61 .22 .08 2, 25 .07 .04 .32 .13	0.35 .90 .44 .04 .10 .09 .52 .30 .71 .72 0 .20 .05 .12 .52 .18 .22 .07 .28 .40 .75 0 0 1.86 .50 .47 0	0.13 1.32 .27 т .10 .10 .57 .42 .62 .05 .03 .61 .05 , 12 .06 .18 .60 .18 .12 т .28 .50 .60 Ω .12 1, 12 .20 .37 0 0	0.278 .990 .320 .076 .110 .072 $\frac{492}{325}$ $.648$ $.040$.070 .522 .012 .260 .042 $.288$ $.542$.178 .230 .170 .165 .355 $.652$ $.090$.105 1.738 .208 .137 .300 .050		

EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY 39

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TABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for cooling purposes, from July 1954 to October 1955*

40 STUDIES OF EVAPORATION

TABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for cooling purposes, from July 1954 to October 195F—*Continued

EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY **41**

TABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for cooling purposes, from July 1954 to October 1955—*Continued

42 STUDIES OF **EVAPORATION**

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TABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for cooling purposes, from July 1954 to October 1955—*Continued

EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY **43**

"ABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for cooling purposes, from July 1954 to October 1955—*Continued

44 STUDIES OF EVAPORATION

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TABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for ccoling purposes, from July 1954 to October 1955—*Continued

EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY 45

TABLE 9.—*Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for cooling purposes from July 1954 to October 1955—*Continued

46 STUDIES OF EVAPORATION

TABLE 9. — *Daily averages of air and water temperatures, wind speed, and flow and temperatures of water used for co-fling purposes, from July 1954 to October 1955 —* Continued

EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY

TABLE 10,—*Mean water temperatures for S-foot layers for each thermal survey, in °C, for time shown*

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48 STUDIES OF EVAPORATION

TABLE 11.—*Lake Colorado City class A evaporation pan data, from July 21, 1954, to October 80, 1955*

* Accumulated evaporation since last observation.

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EFFECT OF ADDITION OF HEAT ON LAKE COLORADO CITY 49

Date 1. 2.. ___ .. ____ 3

5.--------------4..— —— — — - 5. _ — — — —— 6. _ 7-- _ 8. _ ... — _ 9. __ 10.... __ — - _ March 1955 Evapo-ration (inches) $0.222338261523136442772636224440827261313519$ Water tempera-ture (°F) Max 75 82 81 62 53 66 70 75 75 78 62 80 84 78 66 63 55 77 76 54 65 73 77 Min 45 45 46 54 53 40 32 33 35 48 53 53 47 41 61 43 43 46 45 46 35 32 33 39 43 Wind (miles per day) 45 98 68 86 151 96 61 114 114 114 70 62 87 104 168 142 106 63 80 150 103 171 69 178 149 162 143 14 92 184 April 1955 Evapo-ration (inches) 484111348833209983441333143483453343614944433411348834633436345334853485493453534 Water tempera-ture (°F) Mas 70 71 76 79 76 64 70 70 66 83 78 72 73 82 81 87 83 79 84 81 84 89 69 83 83 88 85 82 84 Min 4307084455444645559424555669652586 Wind (miles per day) 1293 59 54 71 36 0 24 33 38 32 78 00 50 30 87 43 44 45 56 97 78 36 56 27 8 43 44 56 97 78 36 56 22 8 May 1955 Evapo-ration (inches) $0.48624530442492861719724363830342753553315676453846924$ Water tempera-ture (°F) Max 89 90 93 11 37 48 31 78 92 92 90 93 87 77 78 90 33 83 83 83 83 83 84 85 87 Min 63 53 63 63 63 61 61 61 62 63 63 60 60 60 67 59 62 60 63 63 59 57 67 54 59 Wind (miles per day) 104 34 83 83 83 83 84 65 83 89 76 75 19 25 31 31 31 31 32 93 94 42 60 27 98 94 54 55 84 42 60 27 98 June 1955 Evapo-ration (inches) 0.41 .45 .25 .50 .35 .40 .41 .15 .23 .40 .35 .48 .51 .48 .49 .37 .46 .51 .40 .34 .43 .45 .52 .63 .63 .37 .76 .51 .47 .64 Water tempera-ture (°F) Max 91 91 91 93 89 90 96 93 70 79 87 90 91 94 94 95 95 93 88 94 93 * 94 96 95 92 95 92 Min 66766644666666666666666768766697666976869 $^{\mathrm{tv}\mathrm{ind}}$ (miles per day) 70 132 100 107 62 105 90 148 132 110 52 100 91 92 78 120 108 82 187 49 61 37 81 99 92 88 97 128 134 120

TABLE 11.—Lake Colorado City class A evaporation pan data, from July 21, 1954, to October 30, 1955—Continued			
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i Accumulated evaporation since last observation.

11..——————— 12————————— 13. _ —— —— _ 14———————— 15. _ —— __ - 16.. — _ 17——————— 18—————————

19- ————— — — 20— ——— — ——— 21-. _ ——— — - 22. ___ — _ ——

23... ___ ... ——— 24 _ 25.———— ———— 26. ————— ——

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 $\label{eq:2.1} \frac{1}{\sqrt{2}}\int_{\mathbb{R}^3}\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\frac{1}{\sqrt{2}}\frac{1}{\sqrt{2}}\frac{1}{\sqrt{2}}\frac{1}{\sqrt{2}}\frac{1}{\sqrt{2}}\frac{1}{\sqrt{2}}$ $\label{eq:2.1} \frac{1}{\sqrt{2}}\int_{0}^{\infty}\frac{1}{\sqrt{2\pi}}\left(\frac{1}{\sqrt{2\pi}}\int_{0}^{\infty}\frac{1}{\sqrt{2\pi}}\left(\frac{1}{\sqrt{2\pi}}\int_{0}^{\infty}\frac{1}{\sqrt{2\pi}}\right)\frac{1}{\sqrt{2\pi}}\right)\frac{d\theta}{\sqrt{2\pi}}\,d\theta.$

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