

Base modified and reprojected from 1:2,000,000 National Atlas of the United States (Streams and Waterbodies, preliminary version, 2003; State Boundaries, 2002; and Cities and Towns, 2002). Waterbodies and cities in Canada from 1:500,000 topographic maps from Natural Resources Canada (Letbridge, 1995) and Department of Energy, Mines and Resources (Fremont, Cypress Hills, Wood Mountain, Willow Bush Lake and Wapeton quadrangles, 1951-1977). Grid created for this map by the authors. Reprojection: Mercator Projection. Datum: NAD 1983. Symbols: CLARK 1985.

LETTERING	FORM	CYPRESS HILLS	WOOD MOUNTAIN	WILLOW BUSH LAKE	WAPETON
CUT BANK	SHOULDER	HARE	SASSON	WOOD POINT	WILLOWTON
DOUGLAS	GREAT FALLS	LEWISTOWN	JORDAN	GLENDEN	WATFORD CITY
					DORNBURG

INDEX TO 1" x 2" TOPOGRAPHIC QUADRANGLE MAPS USED IN PREPARATION OF THIS MAP

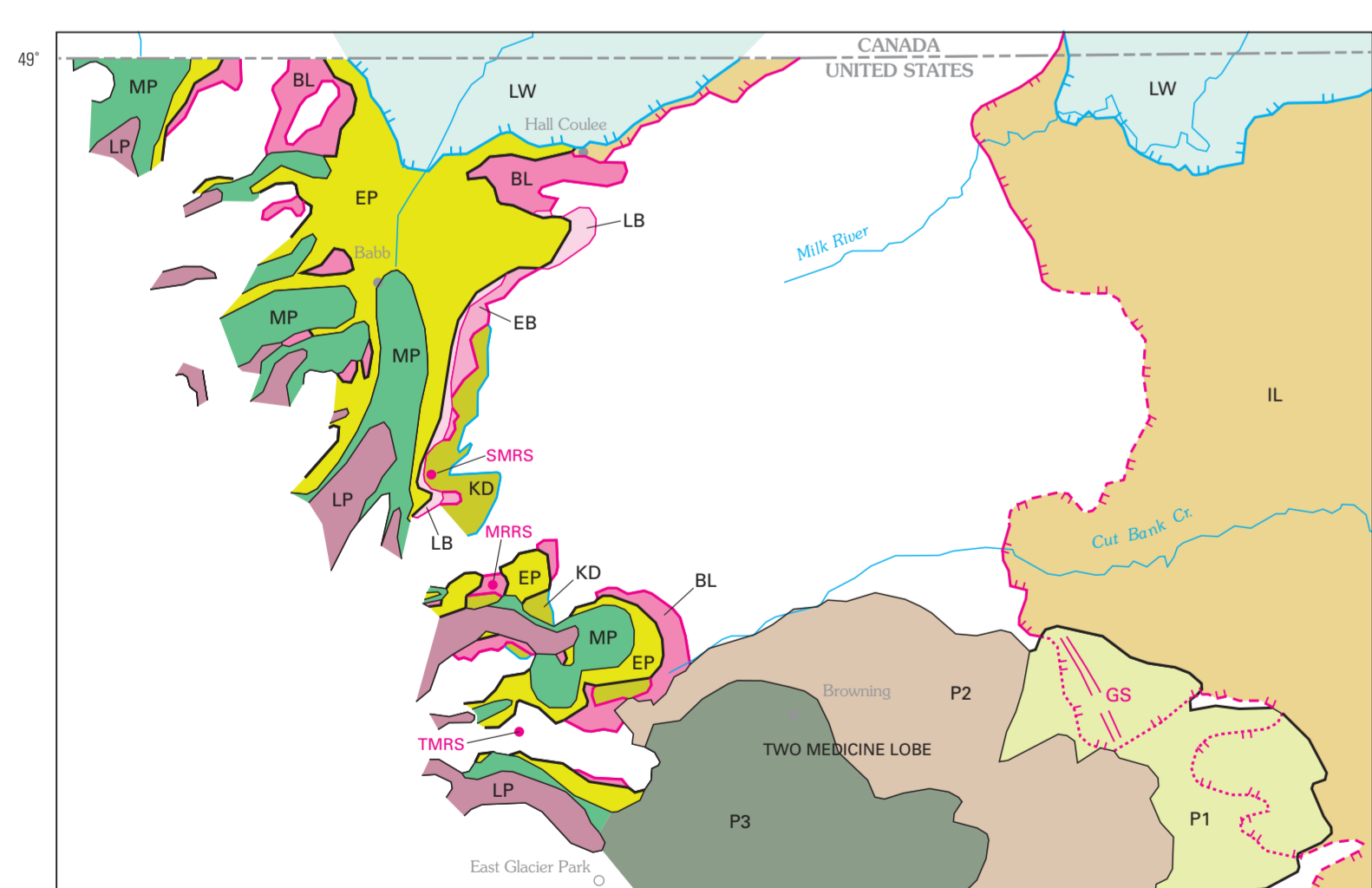


Figure 1. Expanded view of rectangular area near Glacier National Park.

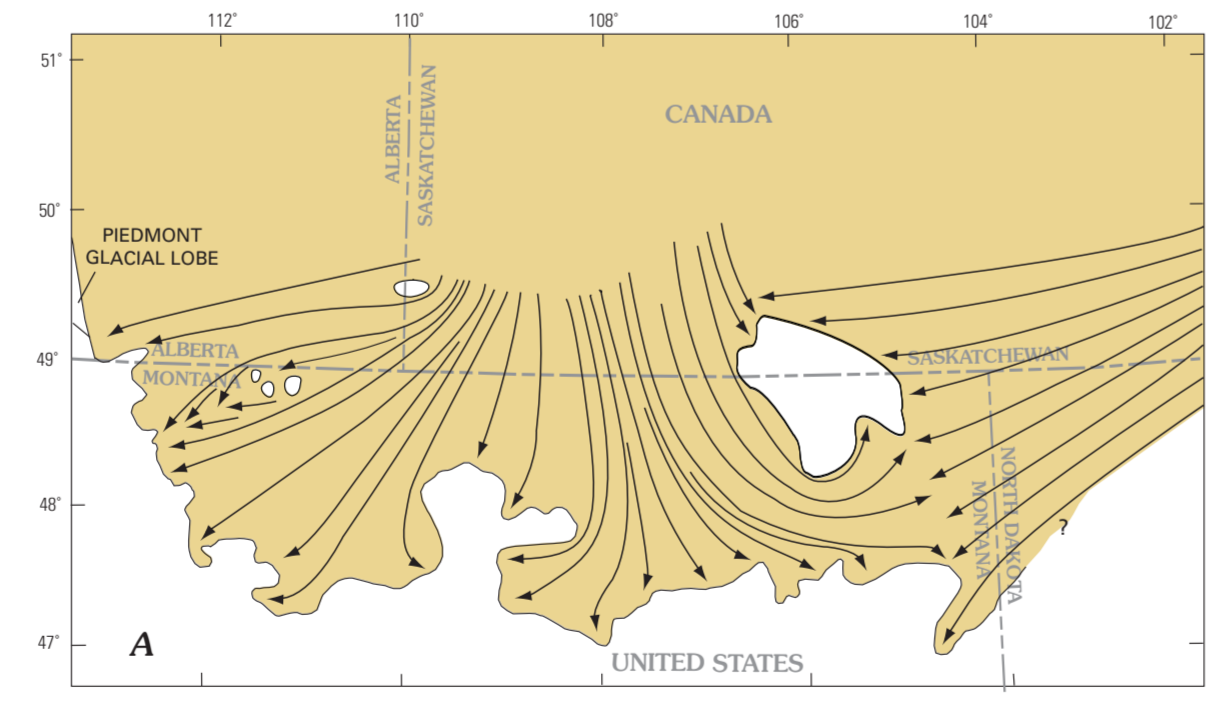
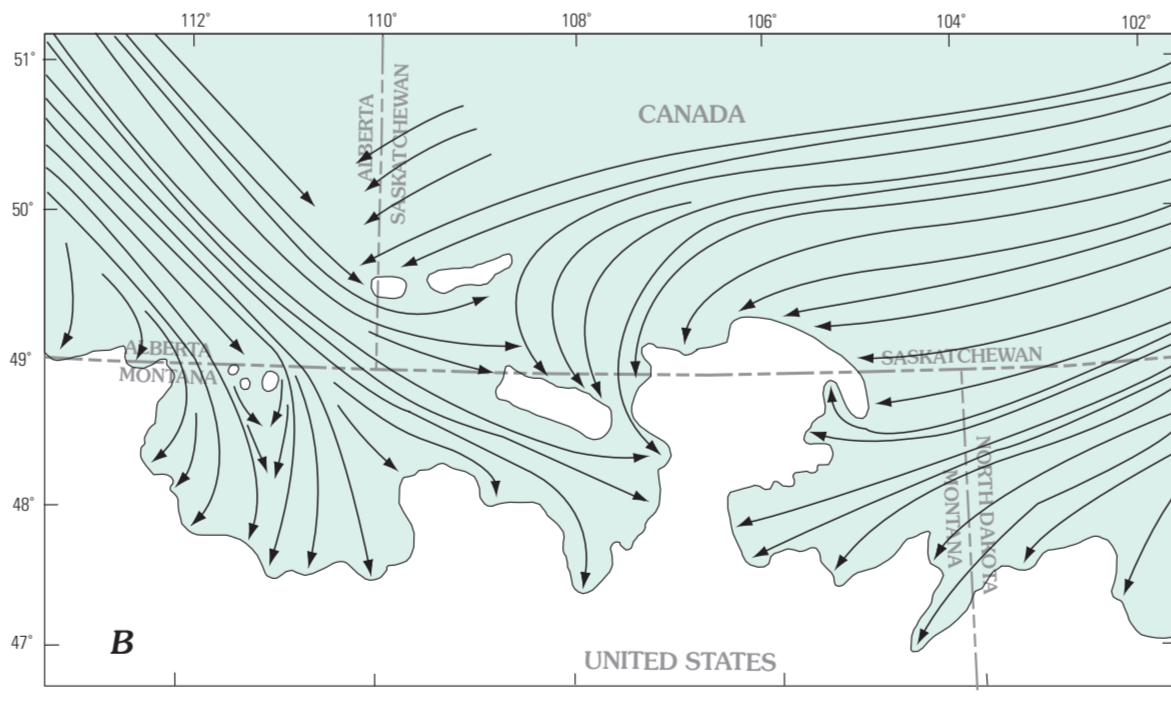
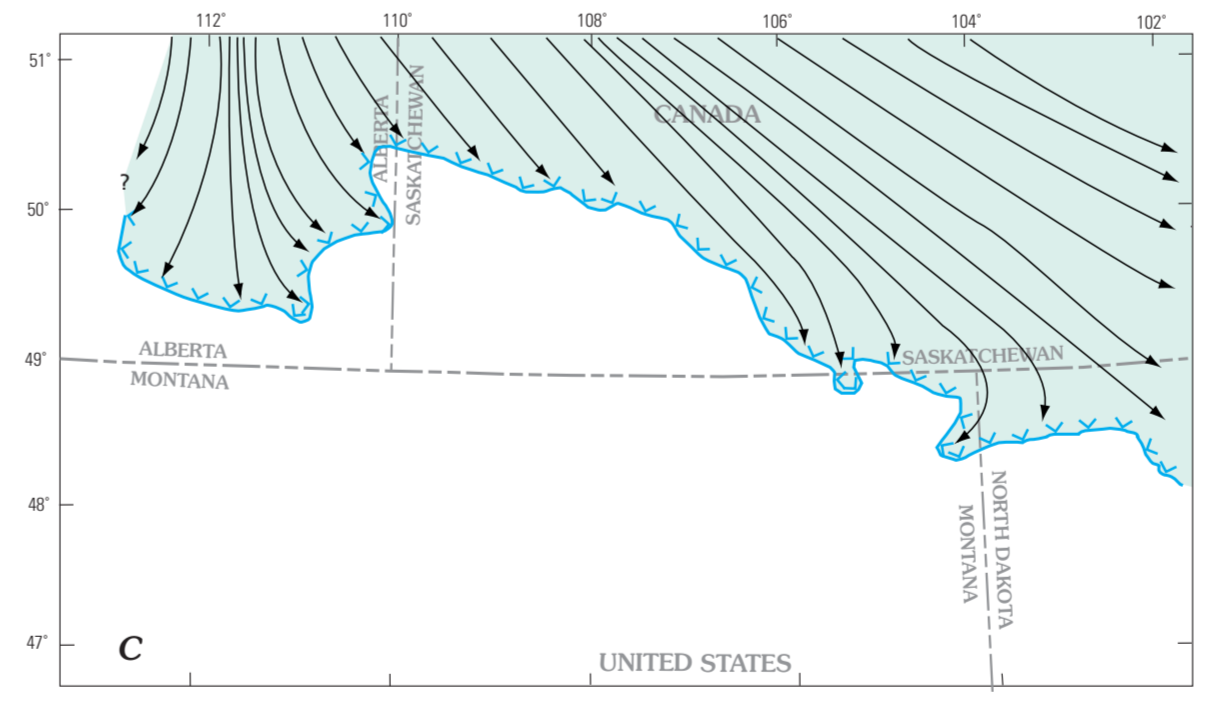


Figure 2. Schematic representation of regional ice-flow patterns during three Laurentide continental glacial maxima. Note that the northern boundary of the map of glacial limits (above) is lat 50° N; the northern boundary in this figure (LA-C) is north of lat 50° N. The west boundary of the map of glacial limits is west of the boundary in this figure. A. Maximum late Wisconsin glacial (~14,000 14C yr B.P.). During maximum late Wisconsin glacial, regional ice flow in southwestern Saskatchewan was southward into Montana. Ice flowed westward between the Sweetgrass Hills ranahtals in Montana and the Cypress Hills (west block) ranahtal in Alberta and Saskatchewan. The central and eastern blocks of the Cypress Hills were glaciated. Westward ice flow in Alberta north and northeast of Glacier National Park was blocked by a large piedmont lobe of coalesced mountain valley-glacier ice in the valleys of the Oldman, Waterton, Belly, and St. Mary Rivers (see pamphlet). Ice from a second, subordinate dispersal center flowed southwestward, west-southwestward, and westward through southwestern Manitoba (east of this map area), southeastern Saskatchewan, and northwestern North Dakota into Montana. The areal distributions and the clast lithologies of the subsurface deposits of the two documented pre-Illinoian Pleistocene continental glaciations in Montana (not mapped) indicate that the regional patterns of ice flow during the maxima of both of those glaciations was similar to that during Illinoian glaciation, and it was very different from the regional flow pattern during maximum late Wisconsin glacial (fig. 1B). B. Maximum late Wisconsin glacial (~20,000 14C yr B.P. or ~23,360



14C yr B.P.). During maximum late Wisconsin glacial, the highest parts of the central and eastern blocks of the Cypress Hills and the Boundary Plateau were not glaciated. The predominant ice-dispersal center was either (1) a Patuxent center south or southwest of Hudson Bay in Canada (Tyrrel, 1914; Johnston, 1933) or a modeled center in that general location (Clark and others, 1996, 1997), (2) a "Hudson dome" in the southwestern part of Hudson Bay (Dyke and others, 1982; Hughes, 1985, 1987), or (3) a "Hudson ice divide" southwest of Hudson Bay (Dyke and Prest, 1987b). The flow was not from a distant dispersal center in Quebec or Labrador (Shultz, 1980, 1982, 1985; Vallette and others, 1999) or from an "ancestral Keewatin" ice divide (Dyke and Prest, 1987b). Ice from a second, subordinate, Keewatin dispersal center west of Hudson Bay was forced far westward in Canada, and coalesced Keewatin ice and "Canadian" ice from the Rocky Mountains (Dyke and Prest, 1987a, b) flowed southwestward from central Alberta into Montana. Southeastward regional flow of ice from central Alberta into Montana occurred only during late Wisconsin glacial. The regional ice flow pattern depicted in figure 1B is incompatible with reconstructions of Laurentide ice sheet flow during maximum late Wisconsin glacial by Dyke and others (1982, Fisher and others (1985), Andrews and Fulton (1987), Dyke and Prest (1987a, b), and Hughes (1987). Later than ~17,000 14C yr B.P. (~19,850 CAL yr B.P.), the "northeastern" dispersal center collapsed and regional ice-flow patterns changed markedly. The ice thinned and stagnated on uplands, and



active ice margins retreated rapidly in lowlands (Fullerton and others, 1995, 2000, in press). The positions of ice domes, ice divides, and ice saddles in Canada shifted (Shultz, 1980, 1982, 1985). The Keewatin center subsequently became the predominant ice-dispersal center for Laurentide glaciation in southwestern Alberta, southeastern Saskatchewan, southwestern Manitoba, northeastern Montana, and northwestern North Dakota. C. Regional readvance maximum (~14,000 14C yr B.P. or ~16,350 CAL yr B.P.) following collapse of a "northeastern" ice-dispersal center. A regional Keewatin-source glacial readvance in Alberta, Saskatchewan, Manitoba, Montana, North Dakota, Minnesota, South Dakota, and Iowa culminated ~14,000 14C yr B.P. (~16,350 CAL yr B.P.) (Fullerton and others, 1995, 2000, in press). The regional ice flow depicted similar to a reconstruction by Dyke and Prest (1987a, b); it is incompatible with a reconstruction of Laurentide ice sheet flow ~14,000 14C yr B.P. by Hughes (1987). Reconstructions of ice flow during maximum late Wisconsin glacial by Dyke and others (1982) and Andrews and Fulton (1987) depicted regional flow at ~14,000 14C yr B.P., not regional ice flow during maximum glacial ~20,000 14C yr B.P. (~23,360 CAL yr B.P.). The CAL ages are 14C ages calculated by calibration of the 14C time scale to absolute years by comparison of 14C ages and 147Sm/147Pu ages of coral samples (see "Introduction" in accompanying pamphlet).

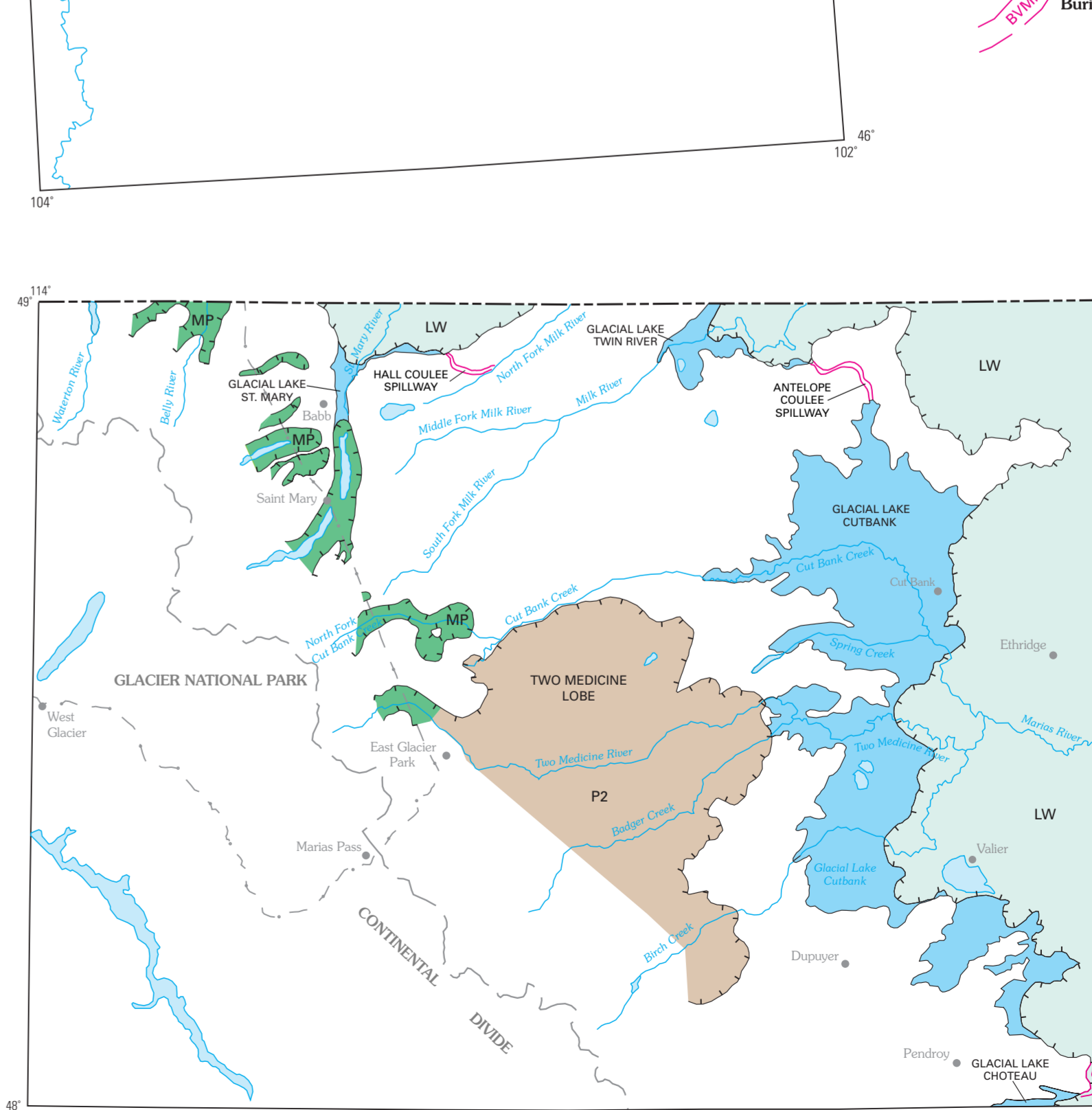


Figure 3. Middle Pinedale and late Wisconsin Laurentide glacial limits in the Cut Bank 1' x 2' quadrangle. LW, late Wisconsin till; MP, middle Pinedale till; P2, Pinedale 2 till. Line decorations on up-glacier side of limit.

- EXPLANATION**
- VALLEY GLACIAL DEPOSITS**
- LP Late Pinedale till
 - MP Middle Pinedale till
 - EP Early Pinedale till
 - LB Late Bull Lake till
 - EB Early Bull Lake till
 - BL Bull Lake till, undivided
 - BL Bull Lake till, undivided
 - KD Kennedy drift
- PIEDMONT GLACIATION GLACIAL DEPOSITS (OUTLET DEPOSITS OF GLACIERS FROM THE NORTHERN MONTANA ICE FIELD)**
- P3 Pinedale 3 till
 - P2 Pinedale 2 till
 - P1 Pinedale 1 till
- CONTINENTAL GLACIAL DEPOSITS**
- LW Late Wisconsin till
 - IL Illinoian till
 - PW Pre-Wisconsin till and glacial erratic boulders
 - PI Pre-Illinoian till (Pleistocene)
 - GB Pre-Illinoian glacial erratic boulders on eroded bedrock (Pleistocene)—Contains no till
- VALLEY GLACIATION**
- Limit of late or middle Pinedale glacial advance (limit of unit LP or MP)
 - Limit of early Pinedale glacial advance or glaciation (limit of unit EP)
 - Limit of late Bull Lake glacial advance or glaciation (limit of unit LB)
 - Limit of early Bull Lake glacial advance or glaciation (limit of unit EB), or limit of Bull Lake till, undivided (unit BL)
 - Mapped limit of pre-Bull Lake Kennedy drift (limit of unit KD)
- PIEDMONT GLACIATION**
- Limit of Pinedale 3 or Pinedale 2 glacial advance (limit of unit P3 or P2)
 - Limit of Pinedale 1 glacial advance or limit of Pinedale glaciation (limit of unit P1)
- CONTINENTAL GLACIATION** [Line decorations on up-glacier side of limit]
- Late Wisconsin**
- Limit of glaciation (limit of unit LW)—Defined by stratigraphy and (or) surface morphology. Dashed where inferred.
 - C2 Limit of C2 regional glacial readvance in southwestern Alberta
 - Limit of regional glacial readvance (see fig. 1C) caused by reorganization of glacial dispersal centers—Defined by stratigraphy and surface morphology. Dashed where inferred.
 - Limit of an ice-margin readvance or position of an ice-margin stillstand—Defined by stratigraphy and (or) surface morphology. Dashed where inferred.
 - Limit of major glaciotectionic deposit or structure—Defined by stratigraphy and (or) surface morphology
- Illinoian**
- Limit of glaciation (limit of unit IL)—Defined by stratigraphy and (or) surface morphology. Dashed where inferred.
 - Limit of a major ice-margin readvance—Defined by surface morphology
 - Limit of an ice-margin readvance or position of an ice-margin stillstand—Defined by surface morphology
- Pre-Wisconsin**
- Limit of till and glacial erratic boulders (unit PW)—Till is not present as far south as the limit of boulders
 - Pre-Illinoian (Pleistocene)
 - Southern limit of known surface exposure of till (unit PI) on north flank of Bearpaw Mountains
 - Buried limit of till (unit PI), overlapped by Illinoian till (unit IL)—Mapped from distribution of exposures of till
 - Limit of glacial erratic boulders on eroded bedrock (unit GB)
 - Stream or shoreline
 - MBS Locality for stratigraphic section—See pamphlet for details. MRS, Milk River Ridge; MBS, Milk River Basin; TMRS, Two Medicine Ridge; SMRS, St. Mary Ridge
 - EB Location of erratic blocks from Canadian Shield
 - LCB Lava Creek B volcanic ash site
 - GS Sag or spillway—CS, Charbonneau Sag; CBS, Caberton Sag; GS, Grandpique Lake Spillway; SSC, Shoshone Sag channels
 - BVMR Buried valley—BVMR, buried valley of ancestral Missouri River; BVYR, buried valley of ancestral Yellowstone River.

MAP SHOWING SPATIAL AND TEMPORAL RELATIONS OF MOUNTAIN AND CONTINENTAL GLACIATIONS ON THE NORTHERN PLAINS, PRIMARILY IN NORTHERN MONTANA AND NORTHWESTERN NORTH DAKOTA

By David S. Fullerton, Roger B. Colton, Charles A. Bush, and Arthur W. Straub 2004



Geologic compilation and interpretation by David S. Fullerton and Roger B. Colton. Digital database by Charles A. Bush. Digital cartography and graphics by Charles A. Bush, Arthur W. Straub, and Diane E. Lane. Edited by Diane E. Lane. Approved for publication June 23, 2004.

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