Geologic Map of the Wrangell-Saint Elias National Park and Preserve, Alaska

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GEOLOGIC NOTE

INTRODUCTION

Wrangell-Saint Elias National Park and Preserve, the largest national park within the U.S. National Park Service system, extends from the northern Pacific Ocean to beyond the eastern Alaska Range into interior Alaska. It features impressively spectacular scenery such as high and craggy mountains, active and ancient volcanoes, expansive ice fields, immense tidewater glaciers, and a myriad of alpine glaciers. The park also includes the famous Kennecott Mine, a world-class copper deposit that was mined from 1911 to 1938 (fig. 1), and remnant ghost town, which is now a National Historic Landmark.

PREVIOUS WORK

Geologic investigations encompassing Wrangell-Saint Elias National Park and Preserve began in 1796, with Dmitriv Tarkhanov, a Russian mining engineer, who unsuccessfully ventured up the Copper River in search of rumored copper (Grinev, 1987, as cited in Hunt, 1996). Lieutenant H.T. Allen (1897) of the U.S. Army made a successful epic summer journey with a limited military crew up the Copper River in 1885, across the Alaska Range, and down the Tanana and Yukon Rivers. Allen’s crew was supported by a prospector named John Bremner and local Eyak and Ahtna native guides whose tribes controlled access into the Copper River basin. Allen witnessed the Ahtnas’ many uses of the native copper. His stories about the copper prompted prospectors to return to this area in search of the rich copper ore in the years following his journey. The region boasts a rich mining and exploration history prior to becoming a park in 1980 (fig. 2). Several U.S. Geological Survey geologists have conducted reconnaissance surveys in the area since Allen’s explorations. This map is the result of their work and is enhanced by more detailed investigations (fig. 3, on map sheet), which began in the late 1950’s and are still continuing. For a better understanding of the processes that have shaped the geology of the park and a history of the geologic investigations in the area, we recommend U.S. Geological Survey Professional Paper 1616, “A Geologic Guide to Wrangell-Saint Elias National Park and Preserve, Alaska”, an exceptionally well illustrated and informative book by Gary R. Winkler, 2000.

BEDROCK TERRANES

Geologically, the park consists of a collage of seven tectonostratigraphic terranes that formed south in the equatorial Pacific Ocean and rafted northward on oceanic plates, eventually accreting to Alaska and the North American continent (fig. 4, on map sheet). Each terrane features a distinct

Figure 1. The Kennecott Mine mill building and railroad cars full of copper ore from the Bonanza and Jumbo Mines of the Kennecott mining district, circa 1920. View to northeast. Photo courtesy of Geoff Bleakley.
stratigraphy and is separated from neighboring terranes by major strike-slip or thrust faults.

The Windy terrane, exposed in the Alaskan interior north of the Denali Fault System and composed largely of Paleozoic metavolcanic and metasedimentary rocks, was the first to arrive; it accreted to the continent in Triassic and early Jurassic time (225–80 Ma).

Next to arrive was the Wrangellia composite terrane (Peninsular, Wrangellia, and Alexander (?) terranes), presently separated from the Windy terrane by the Denali Fault System, a series of strike-slip faults. During its journey north, the Wrangellia terrane merged with the Alexander (?) terrane before docking to the continent in early Cretaceous time (110–85 Ma). The Nutzotin Mountains sequence of Berg and others, 1972, a thick marine flysch of Late Jurassic and Early Cretaceous age, and the short-lived Cretaceous Chisana volcanic arc, formed in and along an oceanic trough next to the north edge of the Wrangellia composite terrane. As the composite terrane progressed north, the Nutzotin Mountain sequence eventually became the glue that attached Wrangellia to the continent.

The Alexander (?) terrane, located on the eastern edge of the park and extending into the neighboring Yukon Territory, Canada, stitched to the Wrangellia terrane by a Pennsylvanian monzonite-granite complex, is composed of early Paleozoic marble, schist, phyllite, and amphibolite.

Wrangellia, the principal terrane in the Wrangellia composite terrane, consists chiefly of a Pennsylvanian-
Permian volcanic arc, probably built on oceanic crust, over lain by a sequence of thick Triassic flood basalts, Triassic limestones, and Jurassic sedimentary rocks. It occupies the bulk of the region with magnificent volcanic edifices and broad glacial valleys.

The Peninsular terrane, buried by the deposits of ancient glacial Lake Atna (Nichols, 1989; Williams and Galloway, 1986) and exposed only locally west of the park, consists chiefly of a Jurassic volcanic arc that overlies the Wrangellia terrane.

Located south of the Wrangellia composite terrane and separated by the Border Ranges Fault, the Southern Margin composite terrane’s spectacular mountains consist of two distinct terranes: the Chugach and Prince William terranes.

The Chugach terrane’s towering peaks are composed largely of another, younger, marine flysch that formed in a deep oceanic trough, which arrived in the latest Cretaceous (about 67 Ma).

Following and separated by the Contact Fault System, the Chugach terrane was joined by the Prince William terrane (about 50 Ma), a Tertiary deep-sea fan and volcanic rock assemblage, home to the ecologically sensitive Copper River Delta.

Finally, arriving about 26 Ma and still actively accreting to the continent, is the Yakutat terrane, a series of Jurassic (?) and Cretaceous to late Tertiary sedimentary rocks that include minor amounts of oceanic volcanic rocks. The Yakutat terrane, host to some of the world’s largest glaciers (fig. 5, on map sheet), is separated from the northern terranes by the Chugach-Saint Elias Fault System and the Fairweather Fault.

QUATERNARY ACTIVITY

The Wrangell volcanic field forms much of the higher elevations of the Wrangell and Saint Elias Mountains. Initial volcanism began about 26 million years ago near the Alaska-Yukon border and progressed northward terminating at the spectacular Quaternary volcanoes of Mount Drum, Mount Wrangell, and Mount Sanford. The volcanic activity is related to the subduction of the Pacific Plate and coincided with the Yakutat terrane’s time of arrival. Subduction and associated volcanic activity largely ended about 200,000 years ago when much of the north movement of the Pacific Plate was then accommodated by right-lateral movement along the Denali Fault System, the Totschunda Fault System, and the Fairweather Fault; all strike-slip faults. As subduction mostly ceased, that part of Alaska south of the Denali-Totschunda Fault Systems, which includes most of Wrangell-Saint Elias National Park and Preserve, was largely coupled to the Pacific Plate. The latest activity to occur on the bounding faults was the shallow Magnitude 7.9 earthquake of November 3, 2002, whose epicenter was on the Denali Fault System northwest of the park. The ground rupture produced along the Denali and the Totschunda Fault Systems stretched into the park, as far southeast as the Nabesna River.

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DESCRIPTION OF MAP UNITS

[Some unit exposures are too small to distinguish for identification. While labeled when possible, all are attributed within the digital database.]

SURFICIAL DEPOSITS

Alluvial deposits

Qa  Alluvium (Holocene)—In active streambeds and lowest terraces along larger streams and active alluvial fans. Chiefly well sorted gravel and sand; minor silt and boulders

Qao  Older alluvium (Holocene and Pleistocene)—In higher terraces along larger streams and in inactive alluvial fans. Predominantly of Pleistocene age, but may include localized Holocene deposits. Chiefly well sorted gravel and sand; minor silt and boulders
Mass wasting deposits

Qu  Slope deposits, undifferentiated (Holocene and Pleistocene)—Colluvium consisting of talus and small landslides, also includes small glacial, rock glacier, and alluvial deposits. Sand, gravel, boulders, diamicton, and rubble

Qcb  Bluff colluvium (Holocene)—Occurs along high steep bluffs of the Copper River and its principal tributaries. Locally includes small areas of the Copper River deposits (Qcr) and the drift of older glaciations (Qg). Unsorted silt, sand, gravel, and boulders

Ql  Landslide deposits (Holocene and Pleistocene)—Slumps, earth-debris flows, block glides, and debris avalanches. Diamicton and rubble

Glacial and glaciolacustrine deposits

Qrg  Rock glacier deposits (Holocene)—Well-defined lobate forms that may include glacier ice cores. Chiefly blocks of rubble

Qag  Drift of Alaskan glaciation (Holocene)—Neoglacial deposits of end, lateral, and ground moraines, kames and outwash proximal to present glaciers; includes small rock glaciers. Diamicton, local gravel, and sand

Qg  Drift of older glaciations (Pleistocene)—Wisconsin glaciation and older end, lateral, and ground moraines, local kames, kame-like deposits, and outwash. Also includes the glaciolacustrine deposits (Qcr) near complex interface with glacial Lake Atna in lower Chitina River valley. Chiefly diamicton, local gravel and sand

Qcr  Copper River deposits (Pleistocene)—Interbedded glaciolacustrine, glacial, and alluvial deposits. Includes deposits of glacial Lake Atna (Nichols, 1989; Williams and Galloway, 1986); historically complex basin built from numerous lake stages in response to multiple glacial advances and retreats. Locally overlain by small fine-grained alluvium and pond deposits (Qfl). Interbedded with the Mount Wrangell valley-fill lavas (Qwwv) along western margin of Mount Wrangell, especially in Chetaslina River valley area. Clay-pebble silt and diamicton, containing local to regionally distributed beds of gravel and sand

Other unconsolidated deposits

Qfl  Fine-grained deposits (Holocene)—In ponds, swamps, lagoons, and along low-gradient streams. Chiefly silt, clay, fine sand, peat, and other organic material

Qms  Mineral spring deposits (Holocene)—Deposited by warm, Na-HCO₃-rich, CO₂-charged, saline spring waters forming the Klawasi Group (Shrub, Upper Klawasi, and Lower Klawasi) mud volcanoes (Motyka and others, 1989) located west of Mount Drum. Also includes one or more inactive deposits west of Capital Mountain; a methane-charged mineral spring found on Copper River’s east bank near Copper Center. Mud and clayey-silt

Qe  Eolian deposits (Holocene)—Dunes on tops of river bluffs, windblown flats, and river bars. Includes informally named Bremner sand dunes, a large barchan field, located at confluence of Bremner and Copper Rivers. Fine-grained sand

Qb  Beach and delta deposits (Holocene)—Wave-deposited, coarse-grained material and driftwood along seacoast; fine-grained sediment in mud flats of Copper River Delta; and marine uplift terraces. Successive spit-building episodes of progressive gradation are preserved in beach sequences along coastal lowland between Cape Yakataga and Bering Glacier. Boulders, gravel, sand, and silt
WRANDELL VOLCANIC FIELD

This Oligocene to Holocene regional province covers more than 10,000 km² in the Wrangell and Saint Elias Mountains built by predominantly large non-explosive shield eruptive centers and summit calderas with a minor sedimentary base deposit. The Wrangell Lava collectively forms the entire volcanic field of effusive flows, tuffs, breccias, lahar deposits, dikes, and domes. The sedimentary nonmarine Frederika Formation was deposited prior to and during early stages of volcanism locally in southern Wrangell Mountains. Individual volcanic descriptions and ages (Richter and others, 1990) can be found in U.S. Geological Survey Bulletin 2072, “Guide to the Volcanoes of the Western Wrangell Mountains, Alaska: Wrangell-St. Elias National Park and Preserve” (Richter and others, 1995)

Individual volcanic eruptive centers

Mount Wrangell volcano (Holocene and Pleistocene, 0.65 Ma and younger)—Enormously bulky massive broad active shield; non-explosive, ice-filled caldera in summit area. Hundreds of voluminous andesitic lavas flowed for 650,000 years, concluding 25,000 years ago; one or two un dated flows might be younger. As the only currently active volcano, present activity is limited to frequent steam plume fumaroles and occasional summit caldera phreatic eruptions

Mount Drum volcano (Pleistocene, 0.6–0.2 Ma)—Constructed during andesitic cone (or shield) building and rhyolite-dacite ring dome emplacement cycles. A late (0.25 Ma) cataclysmic eruption destroyed top and south sides

Mount Sanford volcano (Pleistocene, 0.9–0.5 Ma)—Height and extensive ice cover limit understanding of volcanic structure. Edifice, apparently sourced from three major vents, is an intermingled complex of lava flows and domes, predominantly andesite, including a particularly extensive rhyolite flow

Mount Jarvis volcanoes (Pleistocene, 1.7–1.0 Ma)—North-trending linear succession of centers including at least two summit calderas and eroded andesitic shields; satellite dacite domes and flows

Mount Jarvis volcanoes (Pleistocene and Pliocene, 1.8–0.9 Ma)—Large shield and remnant summit caldera, stripped by erosion, remained to thick, massive intracaldera lavas as serrated topographic highs

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Other volcanic deposits

**Volcaniclastic deposits (Pleistocene)**

**Qvd** Volcanic debris flows—Chetaslina volcanic mass-flow deposit, probably sourced from Mount Drum, exposed along Chetaslina and Copper Rivers; Sanford volcanic debris flow exposed along Sanford and Copper Rivers. Large blocks of lava and varied debris

**Qpf** Pyroclastic flows—The extensive informally named Nadina hot avalanche deposit from Mount Drum’s cataclysmic eruption and a smaller pyroclastic flow (0.3 Ma), possibly derived during an explosive episode from Mount Drum, deposited at Chetaslina River’s two forks. Abundant bombs, scoria, pumice, and ash; lesser andesite and dacite blocks

**Qbc** Cinder cones, flows, and tuffs (Pleistocene)—Olivine basalt and basaltic andesite

**Older volcanic and sedimentary rocks (Pleistocene, Pliocene, or Miocene)**

**QTw** Lava flows and sedimentary rocks—Flat-lying andesite and dacite lava flows; a Tertiary tillite and the fossil-rich conglomeratic sandstone Frederika Formation

**QTwi** Flows of the Ice Fields Plateau—Massive dacite and high-silica andesite lavas capping Ice Fields Plateau east of Nabesna Glacier

**QTwr** Domes and intrusive rocks—Chiefly dacite and rhyolite

**Hypabyssal and plutonic rocks (Pleistocene, Pliocene, Miocene, or Oligocene)**

**QTwd** Dikes—Chiefly andesite and dacite

**Th** Shallow plutons, small stocks, dikes, and sills—Andesite to dacite porphyry, diorite, and quartz diorite. K-Ar dates indicate ages that range from 7.0 to 3.0 Ma, younger than the plutons (Tp) located on north side of Wrangell Mountains. Probably related to Wrangell volcanism, porphyry rocks may be hypabyssal equivalents of the older Wrangell lavas (QTw) and the Chisana Formation (Kc)

**Tg** Granodiorite—Plutonic phase of the Wrangell Lava exposed mainly at Hawkins Glacier head, marginal to Mount Blackburn volcano caldera

**Oldest volcanic rocks (Miocene or Oligocene)**—Sonya Creek (23–20 Ma) and other, unmapped, older Tertiary volcanic centers (26–23 Ma)

**Tw** Lava flows and lahars—Basaltic andesite, andesite, and dacite lava flows; andesitic lahar deposits

**Twr** Rhyolite flow—Obsidian-bearing lava within Sonya Creek caldera

**Twd** Domes—Dacite and rhyolite underlying the lava flows and lahar deposits (Tw) of Sonya Creek volcano

**WINDY TERRANE**

**Sedimentary rocks**

**Mzc** Conglomerate and sandstone (Early Cretaceous (?) and Late Jurassic (?) or older)—Pebble to boulder conglomerate and coarse-grained sandstone. May reflect the Nutzotin Mountains sequence (KJs) of Berg and others (1972) northern shore facies; hornfels alteration along southern contact of the Snag pluton (Tsq)

**Intrusive rocks**

**Tsq** Snag pluton (Tertiary, 50 Ma)—Informally named body of leucocratic biotite quartz monzonite. K-Ar date indicates age of 50 Ma
Metamorphic rocks

Tc Metamorphic complex of Cottonwood Creek of Richter (1976) (Tertiary)—Banded granite and quartz diorite gneiss enclosed in schist and phyllite, most likely part of the Devonian metasedimentary (Dp) rocks. Greenschist to amphibolite metamorphism yields metamorphic age of 20–17 Ma. Complex occurs within Denali Fault System between Nabbesna and Chisana Rivers.

Metamorphic rocks of the Carden Hills (Permian(?), Pennsylvanian, Mississippian, and Devonian(?))

Pv Metavolcanic rocks—Slightly porphyritic meta-andesite and metabasalt; greenschist-facies metamorphism.

Pvc Metavolcaniclastic rocks—Chiefly meta-sandstone, metatuff, and quartzite; greenschist-facies metamorphism.

Dp Metasedimentary rocks (Devonian)—Mainly phyllite, also includes calcareous mica-schist, quartzite, metaconglomerate, and thin marble lenses. Greenschist-facies metamorphism. Contains poorly preserved Devonian fossils.

Ultramafic and associated rocks

Mpu Alpine-type ultramafic rocks and gabbro (Cretaceous(?), Jurassic(?), Triassic, Permian, and Pennsylvanian(?))—Serpentinized peridotite and dunite, locally containing chromitite, leucocratic gabbro, and hornblende gabbro.

WRANGELLIA COMPOSITE TERRANE

PENINSULAR TERRANE
(West of Copper River)

Sedimentary, volcanic, and ultramafic rocks

Jnt Mafic and ultramafic rocks (Middle or Early Jurassic)—Gabbro, pyroxenite, peridotite and dunite. Includes the Nelchina River Gabbronorite, the Tonsina ultramafic, and a mafic sequence of Winkler and others (1980). K-Ar dates indicate ages of 188–171 Ma.

Jt Talkeetna Formation (Early Jurassic and Late Triassic(?))—Dominantly andesitic volcaniclastic rocks, products of Late Triassic and Early Jurassic oceanic volcanic arcs.

WRANGELLIA TERRANE
(Eastern Alaska Range and Northern Wrangell Mountains)

Metamorphic rocks

Pipdm Diorite complex, metamorphic part (Permian(?)) and Pennsylvanian)—Amphibolite-grade metamorphic rocks of mainly quartz-feldspar schist, biotite-hornblende gneiss, and syenite-monzonite gneiss. Jurassic K-Ar ages in probability reflect a metamorphic age. Together with associated late Paleozoic plutons (Pipdp) construct part of the Skolai-Tetelna magmatic arc (Pip, PipS, and PipSc) and Pennsylvanian oceanic crust. Possibly correlative with the gabbro and orthogneiss (Pipgo) in southern Wrangell Mountains.

Intrusive rocks

Kg  Mid-Cretaceous plutons (Cretaceous)—Chiefly discordant plutons of granodiorite and quartz monzonite; also include quartz diorite, diorite, syenodiorite, and trondhjemite phases. K-Ar dates indicate ages of 117–105 Ma. Locally abundant xenoliths

Pdp  Diorite complex, plutonic part (Permian(?) and Pennsylvanian)—Hornblende diorite grading into gneissic diorite. Together with associated metamorphics (Pdpdm), almost certainly represents the Skolai-Tetelna magmatic arc’s (PdpPt, PdpPs, and PdpSc) roots. Possibly correlative with the gabbro and orthogneiss (PdpGo) in southern Wrangell Mountains

Pa  Ahtell pluton (Pennsylvanian)—Quartz monzonite and granodiorite discordant zoned pluton. Together with the monzonite-granite complex (PdpM) in Saint Elias Mountains make part of the Skolai-Tetelna magmatic arc’s (PdpPt, PdpPs, and PdpSc) core. K-Ar data indicates an age of 285 Ma

Sedimentary and volcanic rocks

Kun  Nonmarine sedimentary rocks (Late Cretaceous)—Conglomerate, sandstone, siltstone, and mudstone locally containing abundant plant fossils

Kc  Chisana Formation (Early Cretaceous)—Thick volcanic arc assemblage of marine and subaerial volcanic and volcaniclastic rocks. Younger part consists of andesite and basaltic andesite flows, and volcaniclastic rocks; older part consists of volcaniclastic rocks and locally fossil rich marine argillite and graywacke

KJs  Nutzotin Mountains sequence of Berg and others (1972) (Early Cretaceous and Late Jurassic)—Argillite-graywacke graded beds and minor conglomerate, sandstone, and limestone; thick deep ocean trench fossil-poor marine flysch sequence bordering Wrangellia terrane

Tl  Triassic limestone (Late Triassic)—Locally fossil rich, thin-bedded limestone and interbedded argillite in partially starved marine basin; overlies shallow marine to intertidal fossil-poor cherty massive limestone

Tn  Nikolai Greenstone (Late and(or) Middle Triassic)—Widespread sequence of tholeiitic flood basalt flows, mainly subaerial and characteristically amygdaloidal; subordinate volcaniclastic interbeds and volcanic conglomerate basal sections. Metamorphosed to prehnite-pumpellyite facies although alteration rarely masks primary mineralogy and texture. Contains anomalously high amounts of copper

Mankomen Group (Early Permian and Pennsylvanian)—Marine sedimentary sequence consisting of the upper nonvolcanogenic Eagle Creek Formation (Pe and Pel) and the lower volcanogenic Slana Spur Formation (Pipss). Locally altered and metamorphosed. Correlates with the Skolai Group (PipPs, Ph, Pl, and PipSc) in southern Wrangell Mountains

Pe  Eagle Creek Formation (Early Permian)—Nonvolcanogenic argillite and siltstone deposited in a moderately deep quiet basin. Abundant fossils, cone-in-cone structures, and concretions. Intruded by locally abundant Triassic gabbroic inclusions (Tn) and porphyry dikes and sills (Tp)

Pel  Limestone of Eagle Creek Formation—Fossiliferous thin- to thick-bedded limestone

Pipss  Slana Spur Formation (Early Permian and Pennsylvanian)—Fossil-rich volcanic sandstone, siltstone, conglomerate, and impure bioclastic limestone; shed off the Tetelna volcanic arc (PipT)

PipT  Tetelna Volcanics (Early Permian and Pennsylvanian)—Subaerial and submarine basaltic to dacite lava, lahar flows, tuff deposits, and fine-grained volcaniclastic rocks. Extensive late Paleozoic magmatic arc vestiges form Wrangellia terrane’s basement in northern Wrangell Mountains and eastern
Alaska Range. Correlates to the Skolai Group’s (P*Ps) Station Creek Formation (P*PsC) in southern Wrangell Mountains

WRANGELLIA TERRANE
(Southern Wrangell Mountains and Saint Elias Mountains)

Sedimentary and volcanic rocks

**Ku**  
| Marine sedimentary rocks (Early and Late Cretaceous) | Sandstone and siltstone deposited at shallow to moderate depths; locally abundant marine fossils. Includes the MacColl Ridge, Chititu, Schulze, and Moonshine Creek Formations |

**Kl**  
| Shallow marine sedimentary rocks (Early Cretaceous) | Conglomerate, sandstone, siltstone, and clastic sandy limestone deposited under nearshore, high-energy, transgressive marine conditions; locally includes concretions and abundant fossils. Includes the Kennicott, Kuskulana Pass, and Berg Creek Formations |

**Js**  
| Sedimentary rocks (Late, Middle, or Early Jurassic) | Chiefly spiculite, graywacke, mudstone, and conglomerate deposited in a shallow-marine and nearshore littoral setting. Includes the fossil-rich Lubbe Creek, Nizina Mountain, and Root Glacier Formations |

**Jk**  
| Kotsina Conglomerate | Conglomerate and minor sandstone deposited under rapid local marine circumstances during orogenesis. Fossils absent |

**J*cn**  
| McCarthy Formation (Early Jurassic and Late Triassic) | Spiculite, impure chert and limestone, and shale deposited in a starved marine basin environment. Fossils locally abundant |

**J*n**  
| Chitistone and Nizina Limestones (Late Triassic) | Chiefly limestone and dolomite; minor chert and marble; locally fossil rich. The upper Chitistone and cherty Nizina, deposited in a shallow to moderately deep marine neritic environment, transition into an intertidal-supratidal setting with local sabkha facies full of algal-mat chips, stromatolites, and evaporate relics typical of the dolostone-rich lower Chitistone rocks |

**J*n**  
| Nikolai Greenstone (Late and/or Middle Triassic) | Widespread sequence of tholeiitic flood basalt flows, mainly subaerial and characteristically amygdaloidal, include subordinate volcaniclastic interbeds and volcanic conglomerate basal sections. Metamorphosed to prehnite-pumpellyite facies although alteration rarely masks primary mineralogy and texture. Contains anomalously high amounts of copper |

**P*Ps**  
| Skolai Group, undivided (Early Permian and Pennsylvanian) | Lower late Paleozoic magmatic arc and associated upper sedimentary rocks. Equivalent to the Mankomen Group (Pe, Pel, and P*PsS) and the Tetelna Volcanics (P*PsP) in northern Wrangell Mountains and eastern Alaska Range |

**Ph**  
| Hasen Creek Formation (Early Permian) | Fossil-rich marine argillite, shale, and sandstone; nonvolcanic origin |

**Pl**  
| Limestone of Hasen Creek Formation | Fossil-rich, lenticular, marine limestone; partially equivalent to the Strelna metamorphic marble of Plafker and others (1989) (P*PsPm) |

**P*PsC**  
| Station Creek Formation (Early Permian and Pennsylvanian) | Magmatic arc consists of a younger volcaniclastic member and an older andesitic and basaltic lava flow member, both of submarine origin; locally metamorphosed to greenschist-facies south of Chitina Fault |

Intrusive rocks

**Jc**  
| Chitina Valley batholith (Late Jurassic) | Compositionally diverse quartz monzodiorite, ranging from granite to diorite, generally foliated and locally
gneissic particularly within plutons enclosed by the Strelna Metamorphics of Plafker and others (1989) (Pmsm). K-Ar dates indicate plutonic ages of 157–138 Ma, foliated phase ages are as young as 122 Ma; synchronous with Late Jurassic major regional orogeny perhaps related to subduction along Border Ranges Fault

**Pm** Monzonite-granite complex (Pennsylvanian)—Large composite granite and quartz monzonite pluton containing lesser monzonite, syenite, and gabbro emplaced between Wrangellia and Alexander(?) terranes. K-Ar dates indicate ages of 297–295 Ma. Correlates with the Ahtell pluton (Pa) in northern Wrangell Mountains to construct the Tetelna-Skolai magmatic arc’s (Ptt, PPs, and PPsC) core

**Pg0** Gabbro and orthogneiss (Pennsylvanian)—Interlayered hornblende pyroxene gabbro and monzonitic gneiss; overlain by the Skolai Group (Pps, Ph, Pl, and PPsC) and their metamorphosed equivalent, the Strelna Metamorphics of Plafker and others (1989) (Phsm). Represents, in part, Pennsylvanian oceanic crust correlative with the diorite complex (PIpdm and PIpdp) in northern Wrangell Mountains and eastern Alaskan Range

Metamorphic rocks

**PIpsm** Strelna Metamorphics of Plafker and others (1989) (Jurassic(?), Triassic(?), Permian, Early Pennsylvanian, and Late Mississippian(?))—Dominantly calcareous metasedimentary rocks, schist, phyllite, amphibolite, and marble; lesser quartzite, meta-andesite, and metabasalt in thin to thick thrust sheets containing local pods of the tectonically emplaced ultramafic rock and marble (Pmpm); overlying the gabbro and orthogneiss (Pg0). Includes the Jurassic Uranatina Metaplutonic Complex and the Dadina Schist of Carboniferous and possibly older(?) age exposed near Mount Drum. Metamorphic grade increases closer to the Chitina Valley batholith (Jc) and might represent metamorphosed roots of the Skolai–Tetelna magmatic arc (Ptt, PPs, and PPsC).

**PIpM** Marble—Regionally metamorphosed with poorly preserved and locally placed Permian and Early Pennsylvanian fossils. Partially equivalent to the Permian limestone of the Hasen Creek Formation (Pl)

ALEXANDER(?) TERRANE
(Saint Elias Mountains)

Metamorphic rocks

**Pzk** Kaskawulsh Group of Kindle (1953) (Pennsylvanian, Mississippian, Devonian(?), and older)—A metamorphosed marine sequence of marble, schist, and phyllite; minor meta-tuff, flows, breccia, and meta-conglomerate; probably oceanic crustal progression with carbonate reefs. Poorly preserved Ordovician, Devonian, and possible Mississippian and Pennsylvanian fossils present. Metamorphosed to upper greenschist-facies and locally intensely folded. Coextensive with widely occurring rocks within neighboring Yukon Territory, Canada; potentially representing westernmost known extent of Alexander(?) terrane

SOUTHERN MARGIN COMPOSITE TERRANE

CHUGACH TERRANE

Sedimentary, volcanic, and metamorphic rocks

Valdez Group (Late Cretaceous)—Mainly mudstone and graywacke turbidite beds, carbonaceous phyllite, minor mafic tuff and breccia in a thick regionally
extensive belt of accreted flyschoid deep marine rocks. Metamorphic grade increases southward from lower greenschist-facies to amphibolite-facies.

**Kv**
- **Greenschist-facies flysch**—Flysch metamorphosed to lower greenschist-facies

**Kvv**
- **Metamorphosed basalt and tuff**—Tholeiitic metabasalt and basaltic meta-tuff, local pillow lava and pillow breccia, gabbroic dikes and sills

**Kvm**
- **Amphibolite-facies flysch**—Flysch metamorphosed to amphibolite-facies schist

**Kvg**
- **Gneiss and schist**—Gneiss, banded and migmatite. Includes the amphibolite-facies schist of Mount Augusta of Plafker (written commun., 2000)

**KJm**
- **Mélange units (mid-Cretaceous, Jurassic, and Late Triassic(?))**—Tholeiitic pillowd basalt, chert, tuff, greywacke, argillite, and phyllite-greenstone deposited along a convergent plate margin as early as Late Triassic continuously into middle Cretaceous. Includes the McHugh Complex, a structurally complex and chaotic assemblage of oceanic rocks containing abundant slightly deformed radiolarian chert; and a mélangé-like unit between the McHugh Complex and the Valdez Group (Kv)

**Jlc**
- **Schist of Liberty Creek of Winkler and others (1981) (Jurassic(?))**—Basalt flows, breccia, and tuff; regionally metamorphosed greenschist- and blue-schist-facies Middle Jurassic or older oceanic assemblage protolith

**Intrusive rocks**

**Tgp**
- **Granitic plutons (Eocene, 50 Ma)**—Granite, granodiorite, and diorite intruding Chugach and Prince William terranes; intermediate composition Eocene age dikes not shown on map

**PRINCE WILLIAM TERRANE**

Sedimentary and volcanic rocks

**Orca Group (early Eocene and late Paleocene)**—Weakly metamorphosed flysch and basalt thick submarine fan complex

**Tos**
- **Turbidites**—Mudstone to sandstone, locally fossil rich, submarine gravity flows

**Tov**
- **Basalt**—Mainly tabular bodies of pillowed, massive, and columnar jointed tholeiitic basalt lava flows; locally interbedded minor mudstone and siltstone

**Intrusive rocks**

**Tgp**
- **Granitic plutons (Eocene, 50 Ma)**—Granite, granodiorite, and diorite intruding Chugach and Prince William terranes; intermediate composition Eocene age dikes not shown on map

**YAKUTAT TERRANE**

Sedimentary, volcanic, and metamorphic rocks

**QTy**
- **Yakataga Formation (Holocene(?), Pleistocene, Pliocene, and Miocene)**—Mudstone, siltstone, sandstone, and diamictite in a diverse marine and glacio-marine elastic continental shelf deposit. Abundant mega- and micro-fauna

**Tm**
- **Volcanic and hypabyssal rocks (Miocene and Oligocene)**—Basalt to andesitic plugs, dikes, and flows

**Tps**
- **Poul Creek Formation (early Miocene(?), Oligocene, and late Eocene(?))**—Concretionary marine siltstone, claystone, sandstone, and conglomerate, minor basaltic tuff, and locally glauconitic. Includes part of the clastic Redwood Formation and, as mapped, forms part of Robinson Mountains north of Cape Yakataga
**Tt** Tokun Formation (Eocene)—Shallow marine concretionary siltstone and minor sandstone

**Tk** Kulthieth Formation (Eocene and early Paleocene(?))—Interbedded nonmarine carbonaceous sandstone, siltstone, shale, and minor coal; sparse warm-water marine molluscan megafauna and locally abundant subtropical plant fossils. Includes small occurrences of the siltstone-rich Stillwater Formation east of Steller Glacier, the siltstone of Oily Lake of Plafker (1987), and the basalt of Hubbs Creek of Plafker and others (1987) found in Samovar Hills, a Malaspina Glacier nunatak

**Kyn** Schist of Nunatak Fiord (Cretaceous)—Complexly deformed interbedded mudstone and sandstone, minor volcaniclastic rocks and conglomerate; metamorphosed to greenschist- and epidote-amphibolite facies. Crops up between Boundary and Fairweather Fault Systems and intruded extensively by the Tertiary granitic plutons (Tgp) (Plafker, written commun., 2000)

**Yakutat Group (Cretaceous and Jurassic(?))**—Complexly deformed flysch and mélangé locally recrystallized to greenschist

**Tf** Flysch—Graywacke, siltstone, and argillite

**Kyf** Mélange—Discrete blocks of greenstone, limestone, marble, graywacke, chert, conglomerate, and plutonic rocks, up to several kilometers in greatest dimension, in a sheared matrix of mudstone

**Td** Rhyolite porphyry and granite (Pliocene (?) or Miocene(?))—Discordant hypabyssal plutons and dikes exposed between Mount Augusta and Mount Cook

**REFERENCES**


Grinev, 1897, Zabytaya ekspeditsiya Dmitriya Tarkhanova na Mednuyu reku [The forgotten expedition of Dmitriy Tarkhanov to the Copper River]: Sovetskiaia Etnografiia [Soviet Ethnography], v. 4, p. 88–180. [English translation available in NPS Archives at 240 W. 5th Ave., Anchorage, AK 99501, phone (907) 644–3480]


