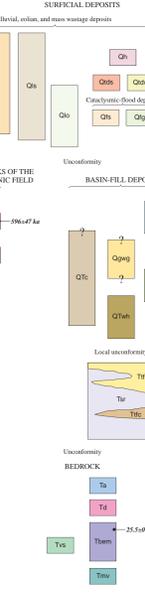


Geologic Map of the Camas Quadrangle, Clark County, Washington, and Multnomah County, Oregon
 By
 Russell C. Evarts and Jim E. O'Connor
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CORRELATION OF MAP UNITS



DESCRIPTION OF MAP UNITS

Artificial fill (Holocene)—Unconsolidated silt, sand, gravel, and crushed rock used for highway and railroad beds and levees.

Alluvium (Holocene and Pleistocene)—Unconsolidated fine sand, silt, and clay that underlies floodplain, islands, and bars of the Columbia River at elevations less than 30 ft (10 m), composed largely of quartz, feldspar, and conspicuous muscovite. Sections exposed in banks, excavations, and drill-core samples commonly show sand-silt couplets up to 30 cm thick; basal sands locally display ripple cross-lamination, planar lamination, and tabular foresets; capping silt and clay beds generally massive and bioturbated, complexly interpreted as deposits of large annual Columbia River freshets before substantial 20th-century river regulation and floodplain diking. Unit includes local diatomaceous beds and layers of organic-rich sediment that probably formed in floodplain marshes, ponds, and lakes, as well as thin (less than 2 cm) tephras from Holocene eruptions of Mount St. Helens. Regionally, most deposits above modern low-water river level (0 ft (0 m) above sea level in map area) are younger than 2000 years old and most deposits forming Groundwater Island are younger than 500 yr B.P. (table 3). Well logs and seismic reflection profiles show that fine-grained silt and clay fill beneath the basalt floodplain, presumably representing river aggradation since the last glacial sea level low stand of about 15,000 yr B.P. (Baker, 2002), locally extends to 300 ft (90 m) below sea level in the map area (Huffsteter, 1984; Hartford and McFarland, 1989; Gates, 1994; Pratt and others, 2001; Rapp, 2005). Near Blue Lake, deposits contain fluviolite concentrated 7.7 ka Murara tephras (Gates, 1994) at 45 ft (14 m) below sea level. Near tributary mouth, unit includes rare lenses of cobble to pebble gravel composed largely of andesite and dacite clasts; also includes local unmappped areas of dredge spoils and other fills, and possibly collin deposits.

Alluvium (Holocene and Pleistocene)—Unconsolidated sand, gravel, and organic-rich mud along rivers and creeks, in seasonally inundated depressions, and along shore of Lacamas Lake. Deposits along the Washougal and Sandy Rivers consist mostly of stratified sand and cobble to pebble gravel, deposits along smaller watercourses are more variable.

Talus deposits (Holocene and Pleistocene)—Unsorted accumulations of angular basalt blocks and scoria below cliffs on south slope of Prune Hill, includes deposits of both natural and anthropogenic origins.

Landslide deposits (Holocene and Pleistocene)—Diamictites of unsorted, angular bedrock and surficial material transported down slope on mass. Chiefly deep-seated, semi-coherent slumps and internally dispersed rockfalls, earthflow, and debris-flow deposits. Many mapped slides head at arcuate scars and exhibit subhorizontal tops, bulbous toes, and hummocks; poorly drained surfaces. Slides at Camas possibly triggered by late Pleistocene cataclysmic flooding.

Deposits derived from Mount Hood volcano (Holocene)—Unconsolidated, stratified sand and gravel and diamictites composed of volcanic debris transported down Sandy River from Mount Hood; development formed by aggradation behind down-river landslide complex in adjacent Washougal quadrangle (R.C. Evarts and J.E. O'Connor, unpub. mapping).

Loess (Pleistocene)—Massive unconsolidated deposits of light-gray to buff, micaceous, quartzofeldspathic colluvial silt and fine sand; commonly contains isolated granules and small pebbles, generally capped with strongly developed red soils. Forms widespread

beds, which are locally reddish to pinkish gray. Gravel clasts are dominantly fresh, porphyritic, pyroxene and hornblende andesites; sand composed mostly of angular to sub-rounded, gray to red, lithic (andesite) fragments, 5–20 percent clear plagioclase fragments, and 2–10 percent equant hornblende crystals. Diamictites are brown to brownish gray, massive to weakly stratified beds, 0.5 to 5 m thick, composed of 30–50 percent angular andesite fragments up to 10 cm diameter in matrix of sand, silt, and clay (Rapp, 2005). Stratified sand and gravel interpreted to represent fluvial aggradation triggered by highly increased sediment supply following Mount Hood eruptions. Diamictites are interpreted as deposits of Mount Hood lahars. Numerous radiocarbon ages (table 3) and regional stratigraphic relations show that these deposits are younger than 1600 yr B.P. and were produced during and after the Timberline and Old Maid eruptive episodes of Craterell (1980); most are older than 1200 yr B.P. and are related to the Timberline period (Rapp, 2005). Old Maid-age sediment forms several channel-fill deposits, as thick as 5 m, in the Sandy River delta and aggraded in the lower Sandy River valley to elevations as high as or higher than the Timberline-age deposits (Rapp, 2005).

Terrace deposits of Sandy River (Holocene and/or Pleistocene)—Poorly exposed, unconsolidated, sandy cobble to boulder gravel forming distinct benches with surface elevation between 170 and 190 ft (48 and 58 m) along Sandy River and Beaver Creek. Consist of subhorizontal gravel sheets, 0.5–5 m thick, locally separated by thin sand lenses; poorly sorted, locally imbricated, and texture varies from compact with sand matrix to loose open-work; clasts subangular to well rounded, include common large (up to 1 m diameter) boulders of hornblende andesite; deposit contains at least one 1-m-thick bed of silty lithic-rich sand inferred to be the distal facies of a Mount Hood lahar. Exhibits weakly developed soil profile and weathering rinds on fine-grained volcanic clasts less than 1 mm thick, suggesting relative youth; absence of cover by cataclysmic (Missoula)-flood deposits indicate that unit postdates flooding; may reflect deposition behind the massive Missoula-flood bar extending southwestward from Brougton Hill that temporarily blocked the Sandy River at its confluence with the Columbia River. Equivalent in part to Estacada Formation of Trimble (1963).

Terrace deposits of lower Washougal River (Holocene and/or Pleistocene)—Unconsolidated sandy gravel and sand underlying small terraces along Washougal River; generally less than 10 m thick. Gravel is poorly sorted, locally imbricated, and texture varies from compact with sand matrix to loose open-work. Clasts subangular to well rounded, derived from Tertiary volcanic and granitic rocks of Cascade Range, Columbia River Basalt Group, and Troadale Formation. Deposits are below 30 ft (15 m) elevation and minimally weathered; inferred to represent aggradation behind coarse-grained Missoula Flood bar that now constricts mouth of Washougal River.

Terrace deposits of Little Washougal River (Holocene and/or Pleistocene)—Unconsolidated gravel and sand flanking Little Washougal River near northeast corner of map area; about 2–5 m thick. Gravel is poorly sorted, composed of subangular to well rounded clasts derived from Tertiary volcanic and granitic rocks of Cascade Range, Columbia River Basalt Group, and Troadale Formation. Minimal soil development. Formed by aggradation behind down-river landslide complex in adjacent Washougal quadrangle (R.C. Evarts and J.E. O'Connor, unpub. mapping).

Loess (Pleistocene)—Massive unconsolidated deposits of light-gray to buff, micaceous, quartzofeldspathic colluvial silt and fine sand; commonly contains isolated granules and small pebbles, generally capped with strongly developed red soils. Forms widespread

mantle on uplands of map area but mapped only where thick (about 3 to 25 m) and extensive enough to obscure underlying units. Overlies 596-ka basaltic andesite of Prune Hill, but probably deposited during several episodes throughout late Quaternary time. Below about 300 ft (90 m) elevation, may include slack-shear cataclysmic-flood deposits (Qc).

Cataclysmic-flood deposits (Pleistocene)—Sediment deposited by colossal glacio-outburst floods caused by repeated failures of ice dams across Clark Fork River that formed Pleistocene Lac Le Sec in western Montana (Breitz, 1925, 1939; Breitz and others, 1956; Trimble, 1963; Allison, 1978; Baker and Bunker, 1985; Wait, 1985, 1994, 1996; Waiter, 1986; O'Connor and Baker, 1992; Benito and O'Connor, 2003). The Missoula floods achieved stages of 400–500 ft (120–150 m) as they spread and slowed over the eastern Portland Basin after exiting the western Columbia River Gorge with velocities of 35 m/s at peak discharge (Benito and O'Connor, 2003) and deposited coarse traction load in series of large bars and plains, the Portland delta of Breitz (1925). Hydraulically damped floodwaters temporarily ponded in Portland Basin and deposited suspended sediment load (Trimble, 1963). Radiocarbon and geochronologic data from outside the map area indicate depositional ages between about 17,000 and 15,000 ¹⁴C years B.P. (Wait, 1985, 1994; Waiter, 1986; Benito and O'Connor, 2003; Clague and others, 2003). Coarse bedload deposits and fine slack-water deposits mapped separately.

Sand and silt facies (Unconsolidated)—Light brown to light gray silt, clay, and fine to medium sand. Up to 20 m thick at low elevations in the region but thin toward upper map edge; extent at 200–350 ft (60–105 m) elevation in the map area. Upper map limit only approximately placed on basis of topographic owing to the difficulty in distinguishing from similar loess and clayey soils that cover most upland surfaces. Most exposures obscure, but fresh exposures show multiple, 0.25-to-1.5-m thick, fining-up sequences of ripple cross-stratified very fine sand grading up to massive bioturbated clay. Sand composed of quartz, feldspar, and conspicuous muscovite, indicative of Columbia River provenance. Coarser sand facies contains abundant dark volcanic rock fragments. Interpreted as slack-water sediment sorted from temporarily ponded floodwaters. May locally include compositionally identical loess.

Gravel facies (Unconsolidated)—Gravel, stratified, bouldery to cobbly gravel and sand deposited in thick sheets over older basin-fill deposits on upland areas flanking the Columbia River, organized into prominent large bar and channel complexes on Mill Plain in the northeast part of map area and much of east Portland, Fairview, and Troutdale south of the river (Allison, 1978; Minervini and others, 2003). Beneath Holocene floodplain deposits (Qc) northeast of Blue Lake, flood gravel occupies buried late glacial channel, channel floor at least 250 ft (70 m) below sea level (Hartford and McFarland, 1989; Pratt and others, 2001). Locally, as along southern edge of the historic Columbia River floodplain and in the Lacamas Lake trough, Missoula-flood deposits from thin (<2 m) and discontinuous mantle on older basin-fill and bedrock units. Numerous active and historic quarry exposures in upland areas reveal more than 15 m of crudely stratified and poorly sorted coarse gravel and sand, commonly deposited in tall (locally >5 m), steeply dipping (up to 35°) facets. Deposits generally fine westward and away from the Columbia River; in some places, contain immense boulders, some exceeding 4 m in long dimension, particularly near Troutdale and about southwest side of Prune Hill. Texturally and compositionally variable; most clasts derived from mafic and granitic rocks of Columbia River Basalt Group and Troadale Formation and Pliocene-Quaternary Cascade Range andesite and other Cascade Range rock types; basaltic andesite, locally, such as southwest of Prune Hill, consists of large, angular-to-subangular blocks eroded from nearby valley walls. Sweeping incision-forms visible in large fresh exposures likely reflect multiple flood events (Benito and O'Connor, 2003). Unit includes local sand accumulations below 200 ft (60 m) elevation, some of them mapped as sand and silt deposits (Qs) by Trimble (1963), that were probably deposited by smaller late-episode floods confined to the narrow Columbia River valley of late Pleistocene time.

Basaltic andesite of the BORING VOLCANIC FIELD

Basaltic andesite of Prune Hill (Pleistocene)—Light to medium-gray, microcrystalline, olivine-phyric, calc-alkaline basaltic andesite (54–65 wt percent SiO₂) that underlies area directly west of Prune Hill, well developed platy and columnar jointing; up to 65 m thick in Fisher Quarry, where complex, jointing patterns and contacts with scoriaceous flow breccia indicate multiple flow lobes. Contains phenocrysts of olivine (5–7 percent, about 0.5 mm, locally to 1 mm across, with minute chromian spinel inclusions), plagioclase (0.1 percent, 0.5–1 mm long), and magnetite (0.1 percent) in trachytic to intergranular, locally microporphyritic groundmass of plagioclase, pyroxene, Fe-Ti oxide, and dark-brown interstitial glass, distinguished by presence of prismatic hypersthene microphenocrysts. Some samples contain controlled and sieved plagioclase phenocrysts about 1 mm across. Well located east of Fisher Quarry in area underlain by scoria deposits (single pattern) and coincident with a strong positive aeromagnetic anomaly (Snyder and others, 1993). Well sorted to poorly sorted scoria beds consist of black to brick-red, variably vesicular clasts as large as 1 m across that are petrographically and chemically similar to associated lava flows; also commonly contain well-rounded pebbles and cobbles of Columbia River Basalt Group and quartzite derived from basin-fill sediments. Normal magnetic polarity (J.T. Hagstrum, written commun., 2000). ⁴⁰Ar/³⁹Ar age of 596±47 ka (table 2) is indistinguishable from ⁴⁰Ar age of 590±50 ka reported by others (1996a).

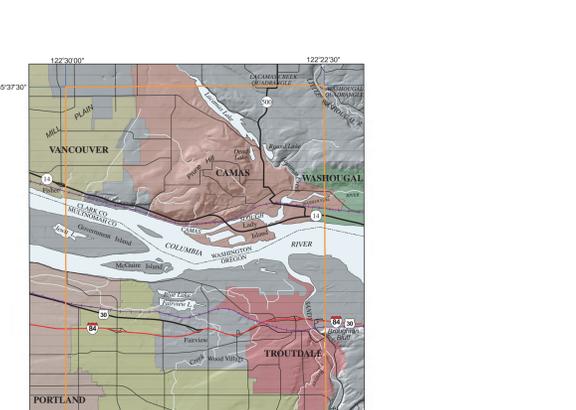
Basaltic andesite of Brougton Hill (Pleistocene)—Light-gray, olivine-phyric, calc-alkaline basaltic andesite flow (52–53 wt percent SiO₂) (R.C. Evarts and R.M. Conroy, unpub. data); flow caps Brougton Hill and has well developed columnar jointing; probably erupted from vent at Chamberlain Hill about 3 km to east. Contains phenocrysts and microphenocrysts of olivine (about 9 percent, mostly 0.3 to 1 mm but a few as large as 3 mm across; with inclusions of chromian spinel locally replaced by ilmenite) in a glassy (volcaniclastic) to intergranular groundmass; sideroclone partly to completely altered to palaginite, which cements sandstone and imparts a distinctive yellowish-brown color to the originally dark-green rock; many beds contain minor admixed mafic and basaltic andesite. Unit locally contains rounded lithic clasts and blackened wood fragments; one exposure displays matrix-poor, talus-like deposit of angular blocks to 2 m across; cemented by clay, zeolites, calcite.

Walters Hill Formation (Pleistocene and/or Pliocene)—Semi-consolidated, deeply weathered, poorly exposed fluvial gravel forming Grant Base. Consists of well rounded to subrounded cobbles and pebbles of volcanic rocks, predominantly porphyritic andesites derived from Cascade Range to the east; matrix is coarse to fine volcaniclastic sand. Age unknown; topographic position suggests unit is younger than Troadale Formation (Tm) and older than other gravel units (Qc, Qm, Qp, Qs) in map area.

Troadale Formation (Pliocene and Miocene)—Semi-consolidated to well consolidated conglomerate and sandstone. Divided into two informal members separated by an unconformity.

Hyaloclastic sandstone member (Pliocene)—Fluvial sedimentary strata distinguished by indurated, coarse sandstone composed of abundant grains of glassy, olivine-plagioclase-phyric basalt and conglomerate that contains olivine-bearing basalt clasts; overlies and intertongues with micaceous, arkosic sandstone, siltstone, and claystone of Columbia River Basalt Group and quartzite derived from basin-fill sediments. Normal magnetic polarity (J.T. Hagstrum, written commun., 2000). ⁴⁰Ar/³⁹Ar age of 596±47 ka (table 2) is indistinguishable from ⁴⁰Ar age of 590±50 ka reported by others (1996a).

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Contact—Dashed where approximately located; short-dashed where inferred; dotted where concealed.

Fault—Dashed where inferred; dotted where concealed; quartered where existence or extent uncertain. Ball and bar on downthrown side. Arrows show relative horizontal movement.

Reverse fault—Dashed where inferred, dotted where concealed. Sawtooth on upper plate.

Strike and dip of beds

Strike and dip of platy parting in lava flows

Basaltic andesite dike

Sample locality for chemical analysis—See table 1

Sample locality for ⁴⁰Ar/³⁹Ar age determination; age in Ma (±1-σ error)—See table 2

Paleoflora locality

Radiocarbon age locality—See table 3

Cinder cone

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