

**In cooperation with the Harris-Galveston Subsidence District, City of Houston,  
Fort Bend Subsidence District, and Lone Star Groundwater Conservation District**

# **Water-Level Altitudes 2010 and Water-Level Changes in the Chicot, Evangeline, and Jasper Aquifers and Compaction 1973–2009 in the Chicot and Evangeline Aquifers, Houston- Galveston Region, Texas**

Scientific Investigations Map 3138

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By Mark C. Kasmarek, Michaela R. Johnson, and Jason K. Ramage

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Scientific Investigations Map 3138

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## Conversion Factors and Datums

### Inch/Pound to SI

Multiply	By	To obtain
Length		
inch (in.)	2.54	centimeter (cm)
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
Area		
square mile (mi <sup>2</sup> )	2.590	square kilometer (km <sup>2</sup> )

### Datums

Vertical coordinate information is referenced to the National Geodetic Vertical Datum of 1929 (NGVD 29) or the North American Vertical Datum of 1988 (NAVD 88). Altitude, as used in this report, refers to distance above the vertical datum.

Horizontal coordinate information is referenced to the North American Datum of 1927 (NAD 27).

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# Water-Level Altitudes 2010 and Water-Level Changes in the Chicot, Evangeline, and Jasper Aquifers and Compaction 1973–2009 in the Chicot and Evangeline Aquifers, Houston-Galveston Region, Texas

By Mark C. Kasmarek, Michaela R. Johnson, and Jason K. Ramage

## Abstract

Most of the subsidence in the Houston-Galveston region has occurred as a direct result of groundwater withdrawals for municipal supply, industrial use, and irrigation that depressed and dewatered the Chicot and Evangeline aquifers causing compaction of the clay layers of the aquifer sediments. This report, prepared by the U.S. Geological Survey, in cooperation with the Harris-Galveston Subsidence District, City of Houston, Fort Bend Subsidence District, and Lone Star Groundwater Conservation District, is one in an annual series of reports depicting water-level altitudes and water-level changes in the Chicot, Evangeline, and Jasper aquifers and compaction in the Chicot and Evangeline aquifers in the Houston-Galveston region. The report contains maps showing 2010 water-level altitudes for the Chicot, Evangeline, and Jasper aquifers, respectively; maps showing 1-year (2009–10) water-level-altitude changes for each aquifer; maps showing 5-year (2005–10) water-level-altitude changes for each aquifer; maps showing long-term (1990–2010 and 1977–2010) water-level-altitude changes for the Chicot and Evangeline aquifers; a map showing long-term (2000–10) water-level-altitude change for the Jasper aquifer; a map showing locations of borehole extensometer sites; and graphs showing measured compaction of subsurface material at the extensometers from 1973, or later, through 2009. Tables listing the data used to construct each aquifer-data map and the compaction graphs are included.

Water levels in the Chicot, Evangeline, and Jasper aquifers were measured during December 2009–March 2010. In 2010, water-level-altitude contours for the Chicot aquifer ranged from 200 feet below National Geodetic Vertical Datum of 1929 or North American Vertical Datum of 1988 (hereinafter, datum) in a small area in southwestern Harris County to 200 feet above datum in central to southwestern Montgomery County. Water-level-altitude changes in the Chicot aquifer ranged from a 49-foot decline to a 67-foot rise (2009–10), from a 25-foot decline to a 35-foot rise (2005–10), from a 40-foot decline to an 80-foot rise (1990–2010), and from a

140-foot decline to a 200-foot rise (1977–2010). In 2010, water-level-altitude contours for the Evangeline aquifer ranged from 300 feet below datum in north-central Harris County to 200 feet above datum at the boundary of Waller, Montgomery, and Grimes Counties. Water-level-altitude changes in the Evangeline aquifer ranged from a 58-foot decline to a 69-foot rise (2009–10), from an 80-foot decline to an 80-foot rise (2005–10), from a 200-foot decline to a 220-foot rise (1990–2010), and from a 320-foot decline to a 220-foot rise (1977–2010). In 2010, water-level-altitude contours for the Jasper aquifer ranged from 200 feet below datum in south-central Montgomery County to 250 feet above datum in eastern-central Grimes County. Water-level-altitude changes in the Jasper aquifer ranged from a 39-foot decline to a 39-foot rise (2009–10), from a 110-foot decline to no change (2005–10), and from a 180-foot decline to no change (2000–10).

Compaction of subsurface materials (mostly in the clay layers) composing the Chicot and Evangeline aquifers was recorded continuously at 13 borehole extensometers at 11 sites. For the period of record beginning in 1973, or later, and ending in December 2009, cumulative clay compaction data measured by 12 extensometers ranged from 0.088 foot at the Texas City-Moses Lake site to 3.559 foot at the Addicks site. The rate of compaction varies from site to site because of differences in groundwater withdrawals near each site and differences among sites in the clay-to-sand ratio in the subsurface materials. Therefore, it is not possible to extrapolate or infer a rate of clay compaction for an area based on the rate of compaction measured at a nearby extensometer.

## Introduction

The Houston-Galveston region comprises Harris, Galveston, Fort Bend, Waller, and Montgomery Counties and adjacent parts of Brazoria, Grimes, Walker, San Jacinto, Liberty, and Chambers Counties (fig. 1) and represents one of the largest areas of subsidence in the United States. Allen

## 2 Water-Level Altitudes 2010 and Water-Level Changes in the Chicot, Evangeline, and Jasper Aquifers



**Figure 1.** Location of Houston-Galveston region study area, Texas.

(1969) describes ground surface displacement (subsidence) as the last step of a variety of subsurface displacement mechanisms that (among others) includes compaction of sediments by loading, drainage, vibration, and hydrocompaction. By 1979, as much as 10 feet (ft) of subsidence had occurred in the Houston-Galveston region, and approximately 3,200 square miles (mi<sup>2</sup>) of the 11,000-mi<sup>2</sup> geographic area had subsided more than 1 ft (Coplin and Galloway, 1999, p. 40). Comparing land-surface altitudes for 1915–17 to those for 2001, Kasmarek, Gabrysch, and Johnson (2009, sheet 2) determined that as much as 13 ft of subsidence has occurred in southeastern Harris County. Groundwater withdrawn from the Chicot, Evangeline, and Jasper aquifers has been the primary source of water for municipal supply and for commercial and industrial uses in the Houston-Galveston region since the early 1900s (Kasmarek and Robinson, 2004). Land-surface subsidence caused by fluid withdrawals was first documented in the Houston area in conjunction with the Goose Creek oil field in southeastern Harris County (Pratt and Johnson, 1926). Most of the subsidence in the Houston-Galveston region has occurred as a direct result of groundwater withdrawals for municipal supply, industrial use, and irrigation that depressed and dewatered the Chicot and Evangeline aquifers causing compaction of the clay layers of the aquifer sediments (Winslow and Doyle, 1954; Winslow and Wood, 1959; Gabrysch and Bonnet, 1975; Gabrysch, 1984; Holzer and Bluntzer 1984; Kasmarek, Houston, and Ramage, 2009).

Subsidence is of particular concern in low-lying coastal areas such as the Houston-Galveston region. Land subsidence in the region has increased the frequency and severity of flooding (Coplin and Galloway, 1999). Low-pressure weather systems such as tropical storms and hurricanes cause high tides and high rates of precipitation, and subsidence exacerbates the effects of storm surge and impedes stormwater runoff by creating topographically low areas where water accumulates. Subsidence has shifted the coastline in parts of the Houston-Galveston region and changed the distribution of wetlands and aquatic vegetation (Coplin and Galloway, 1999). To address the issue of subsidence and its consequences, the Texas State Legislature in 1975 authorized the establishment of the Harris-Galveston Subsidence District (HGSD) to regulate and reduce groundwater withdrawals in Harris and Galveston Counties. In cooperation with the HGSD, the USGS has monitored water levels in the Chicot and Evangeline aquifers and clay compaction in Harris and Galveston Counties since 1976. The USGS has published annual reports of water-level altitudes and water-level changes for the Chicot and Evangeline aquifers in the Houston-Galveston region beginning with the 1977 water levels and annual reports of same for the Fort Bend subregion (Fort Bend County and adjacent areas) beginning with the 1990 water levels. The USGS published its first annual reports of water-level altitudes and water-level changes for the Jasper aquifer in the greater Houston area (primarily Montgomery County) beginning in 2000. Compaction in the Houston-Galveston region has been presented in USGS reports of annual water-level altitudes and

water-level changes since 1981 (compaction for 1973–81) and periodically was reported on separately by USGS authors since 1954 (for example, Winslow and Doyle, 1954; Gabrysch, 1984).

Subsequent to establishing the HGSD, the Texas State Legislature established an additional subsidence district (Fort Bend Subsidence District [FBSD]) and a groundwater conservation district (Lone Star Groundwater Conservation District [LSGCD]) in the Houston-Galveston region to provide for the regulation of groundwater in fast-growing, inland areas (fig. 1). The FBSD was established in 1989 and has jurisdiction throughout Fort Bend County. The FBSD is divided into area A, which includes the Richmond-Rosenberg sub-area, and area B. The LSGCD was established in 2001 and has jurisdiction throughout Montgomery County (fig. 1). The primary purpose of the HGSD and FBSD is to regulate groundwater withdrawal to prevent subsidence that contributes to flooding (Harris-Galveston Subsidence District, 2010; Fort Bend Subsidence District, 2010). The directive of the LSGCD is to conserve, protect, and enhance the groundwater resources of Montgomery County (Lone Star Groundwater Conservation District, 2010). Regulations to gradually change the source of water from groundwater to surface-water supplies are being phased in; the historical, current (2010), and future groundwater management plans of each district are available from their respective websites (Fort Bend Subsidence District, 2010; Harris-Galveston Subsidence District, 2010; Lone Star Groundwater Conservation District, 2010). In the seven other counties of the Houston-Galveston region (fig. 1, inset), groundwater withdrawals remain unregulated.

In 1976 the HGSD began implementing their first groundwater regulatory plan (Harris-Galveston Subsidence District, 2010). An extensive well monitoring network was established by 1977, and water-level data were collected and used to create the first published water-level-altitude maps of the Chicot and Evangeline aquifers in the Houston-Galveston area (Gabrysch, 1979). The FBSD adopted its groundwater management plan in 1990 (Fort Bend Subsidence District, 2010). After a comprehensive well monitoring network was established in Fort Bend, Harris, Brazoria, and Waller Counties in 1989 and 1990, the first water-level altitudes were published for the Chicot and Evangeline aquifers in 1991 (Barbie and others, 1991), and when updated well information became available, the current water-level altitudes were updated in 1997 (Kasmarek, 1997). Similarly, after the creation of the LSGCD in 2001, the USGS first published a water-level-altitude map for the Jasper aquifer in the Houston area (primarily Montgomery County) (Coplin, 2001). In 2004 and again in 2007, as additional wells with reliable water-level-measurement data were inventoried, revised water-level-altitude maps for the Jasper aquifer were prepared (Kasmarek and Lanning-Rush, 2004; Kasmarek and Houston, 2005, 2007). Compared to the 2001 and 2004 maps, the 2007 water-level map is the most comprehensive; the number of wells with reliable water-level data for the Jasper aquifer in the Houston area was sufficient in 2007 to produce a detailed map.

## Purpose and Scope

This report is one in an annual series of reports that depicts water-level altitudes and water-level changes in the Chicot, Evangeline, and Jasper aquifers and compaction in the Chicot and Evangeline aquifers in the Houston-Galveston region. The report also describes the hydrogeology of the study area and provides an overview of the mechanism of clay compaction and land-surface subsidence.

This report contains maps (sheets 1–14) depicting water-level altitudes for 2010 for each of the three aquifers; maps depicting 1-year (2009–10) water-level changes for each aquifer; maps depicting 5-year (2005–10) water-level changes for each aquifer; maps depicting long-term (1990–2010 and 1977–2010) water-level changes for the Chicot and Evangeline aquifers, respectively; and a map depicting long-term (2000–10) water-level change for the Jasper aquifer. Water-level-altitude maps for 1977, 1990, and 2000 are the first maps published as part of the annual series for HGSD, FBSD, and LSGCD, respectively.

In addition to maps depicting water-level altitudes and changes in the Chicot, Evangeline, and Jasper aquifers, this report also contains graphs showing measured compaction of subsurface material in the Chicot and Evangeline aquifers for 12 borehole extensometers at 11 sites in Harris and Galveston Counties from 1973, or later, through 2009. Tables listing the data used to construct each of the maps and the compaction graphs also are included as well as a brief description of the methods used for map construction. Additionally, information about the hydrogeology of the Gulf Coast aquifer system is presented, and a brief description of the regional-scale maps of water-level altitudes and changes for the three aquifers is given.

In the Houston-Galveston region, water-level-altitude contours represent 2010 regional-scale depictions of the water levels in the Chicot, Evangeline, and Jasper aquifers. The areal extent and location of these contours represent the combined effect of groundwater withdrawals at all groundwater wells in the study area. Similarly, water-level-change contours represent regional-scale depictions of water-level change during selected periods for each aquifer. Delineated areas showing contours of water-level rise or decline represent water-level changes in the aquifers caused by spatial and temporal changes in groundwater withdrawals. In areas of water-level rise, clay compaction is minimal or nonexistent, but in areas experiencing appreciable water-level declines, as dewatering and depressurizing of the clay layers occurs, the clay layers are more likely to compact causing associated land-surface-altitude loss in the affected areas.

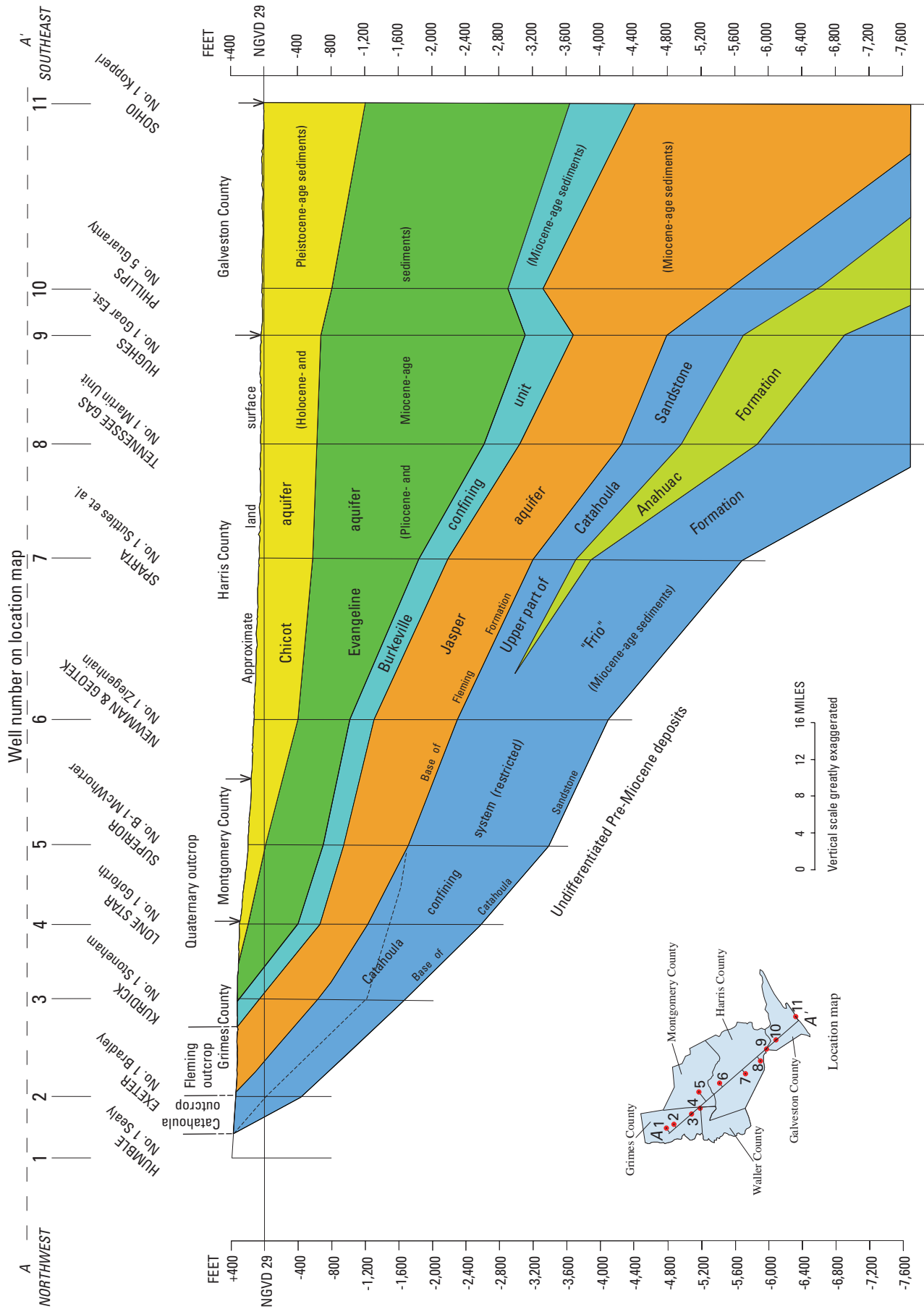
## Hydrogeology of the Study Area

The three primary aquifers in the Gulf Coast aquifer system are the Chicot, Evangeline, and Jasper (figs. 2–4), which are composed of laterally discontinuous deposits of gravel, sand, silt, and clay. The uppermost Chicot aquifer consists

of Pleistocene- and Holocene-age sediments; the Evangeline aquifer consists of Miocene- and Pliocene-age sediments; and the lowermost Jasper aquifer consists of Miocene-age sediments (Baker, 1979, 1986). Through geologic time, geologic and hydrologic processes created accretionary-sediment wedges (stacked sequences of sediments) more than 7,600 ft thick at the coast (fig. 2) (Chowdhury and Turco, 2006). The sediments composing the Gulf Coast aquifer system were deposited by fluvial-deltaic processes and subsequently were eroded and re-deposited (reworked) by large episodic changes in sea level that occurred as a result of oscillations between glacial and interglacial climate conditions (Lambeck and others, 2002). The Gulf Coast aquifer system comprises hydrogeologic units that dip and thicken from northwest to southeast; the aquifers thus crop out in bands inland from and approximately parallel to the coast and become progressively more deeply buried and confined toward the coast (fig. 2). The Burkeville confining unit separates the Evangeline and Jasper aquifers and restricts groundwater flow between the two aquifers. There is no confining unit between the Chicot and Evangeline aquifers; therefore the aquifers are hydraulically connected, allowing groundwater flow between the aquifers (fig. 2). Because of this hydraulic connection, changes in hydraulic head in the Evangeline aquifer affect water levels in the Chicot aquifer. Evidence of this can be seen by comparing the long-term change maps (1977–2010, sheets 5 and 10) and observing that areas where water levels have declined or risen in the Chicot aquifer are approximately coincident with areas where water levels have declined or risen, respectively, in the Evangeline aquifer. The Chicot aquifer can be differentiated from the geologically similar Evangeline aquifer on the basis of hydraulic conductivity (Carr and others, 1985, p. 10) and where each aquifer outcrops—the Chicot outcrops closer to the coast compared to the Evangeline. The Jasper aquifer outcrops inland from the Evangeline aquifer. The Jasper can be differentiated from the Evangeline aquifer on the basis of water levels, which are shallower (closer to land surface) in the Jasper aquifer compared to those in the Evangeline aquifer. In the downdip parts of the aquifer system, the Jasper can be differentiated from the Evangeline aquifer on the basis of stratigraphic position in relation to the Burkeville confining unit (figs. 2–4).

The hydrogeologic cross section A–A' (fig. 2) extends through the Houston-Galveston region from northwestern Grimes County, southeastward through Montgomery and Harris Counties, terminating at the coast in Galveston County, and shows the three aquifers thickening and dipping toward the coast. Comparing cross section A–A' (fig. 2) to cross sections B–B' (fig. 3) in Fort Bend County and C–C' (fig. 4) in Montgomery and Harris Counties, the thicknesses of the three aquifers similarly increase and thicken toward the coast in these counties. In central Harris and southern Montgomery Counties, the sediments of the updip Chicot and Evangeline aquifers become progressively less thick (fig. 2), and in northern Montgomery County, the sediments of the Chicot aquifer are effectively nonexistent for groundwater withdrawal (fig. 4).





**Figure 2.** Hydrogeologic section of the Gulf Coast aquifer system in Harris County and adjacent counties, Texas (modified from Baker, 1979, fig. 4).

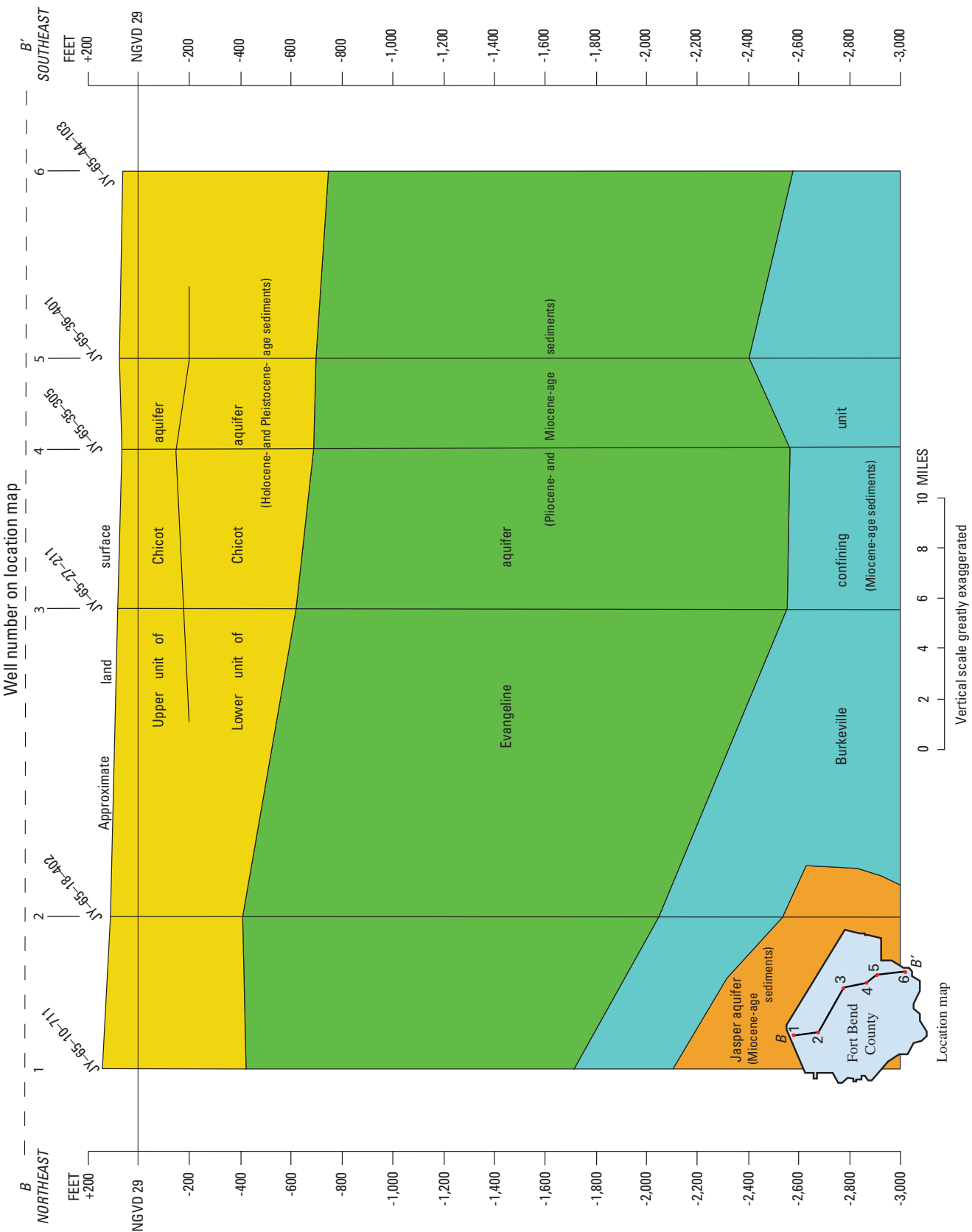
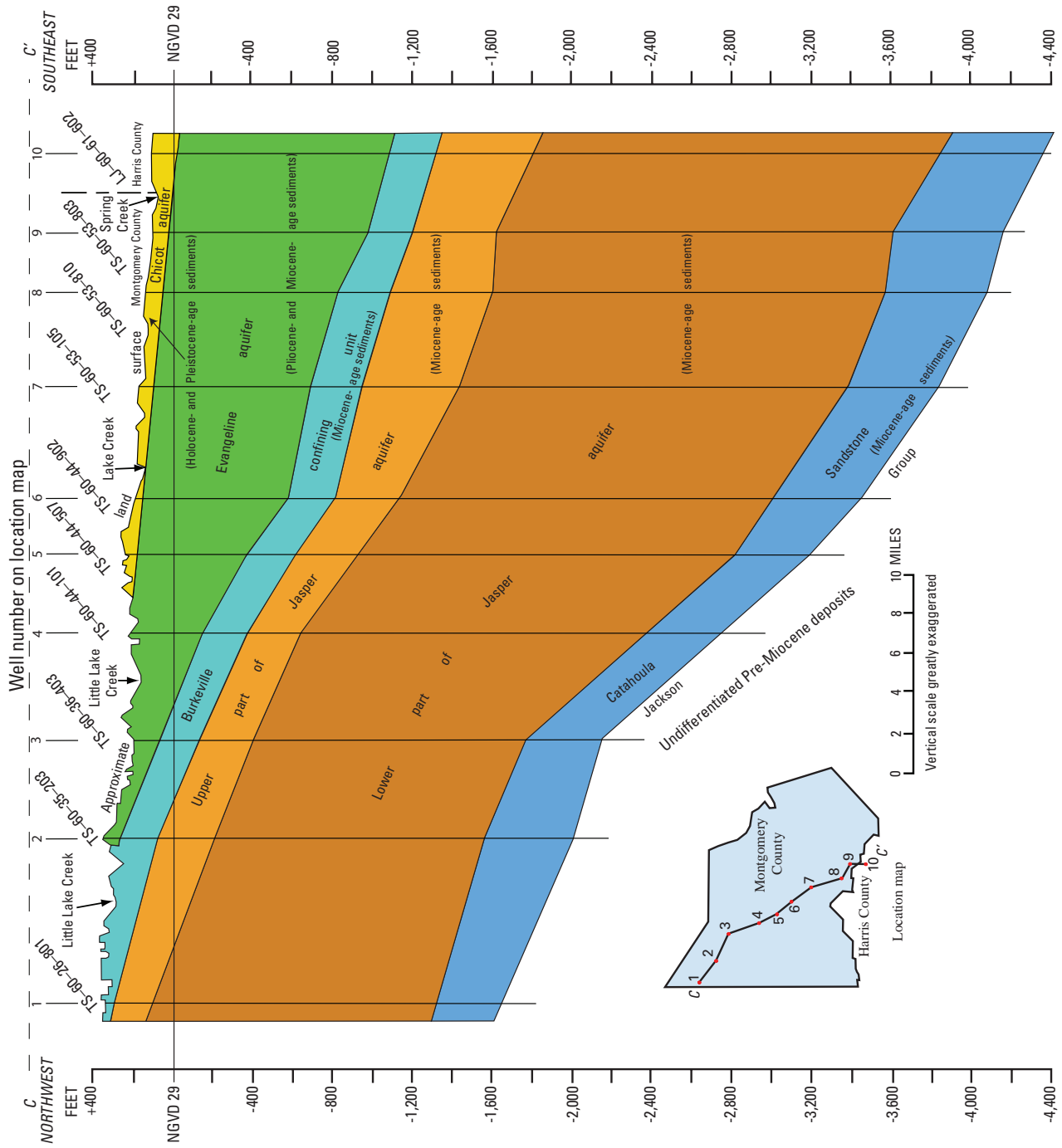


Figure 3. Hydrogeologic section of the Gulf Coast aquifer system in Fort Bend County, Texas (modified from Wesselman, 1972, fig. 30).





**Figure 4.** Hydrogeologic section of the Gulf Coast aquifer system in Montgomery and Harris Counties, Texas (modified from Popkin, 1971, fig. 29).

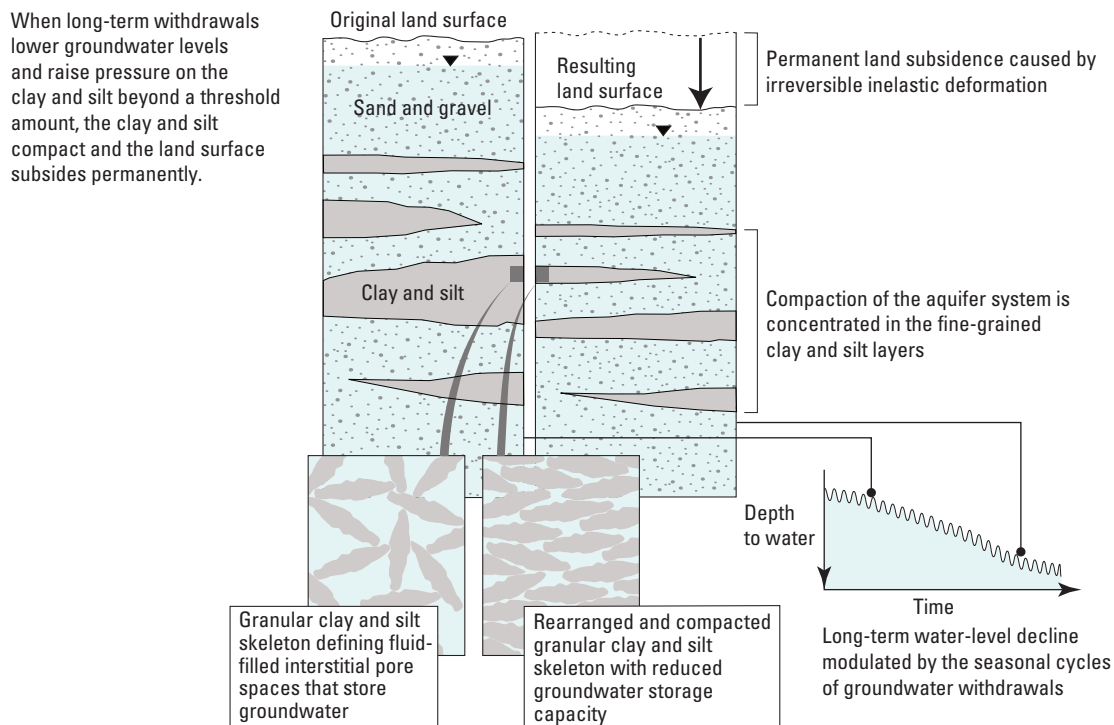
Water in the Chicot, Evangeline, and Jasper aquifers is fresh (less than 1,000 milligrams per liter [mg/L] dissolved-solids concentration) in the Houston-Galveston region, but becomes more saline in the downdip and more deeply buried parts of the aquifers near the coast (Baker, 1979). In the natural groundwater-flow system, water recharges the aquifers in the unconfined outcrop areas, moves downward and coastward, and discharges upward as diffuse upward leakage in the confined downdip areas (Kasmarek, Houston, and Ramage, 2009). Water that does not discharge to streams flows to intermediate and deep zones of the system southeastward of the outcrop areas, and there the water is discharged by wells or by upward leakage in topographically low areas near the coast (Kasmarek and Robinson, 2004).

Water with dissolved-solids concentrations greater than 1,000 mg/L is present in all three aquifers near the coast and at depth. This more-saline water causes less-dense freshwater that has not been captured and discharged by wells to be redirected upward as diffuse upward leakage to shallow zones of the confined downdip areas of the aquifer system and is ultimately discharged to coastal brackish water bodies (Kasmarek and Robinson, 2004).

## Subsidence and Compaction Processes

Subsidence can occur as a result of potentiometric-surface declines in unconsolidated confined aquifers

(Galloway and others, 1999; Kasmarek, Houston, and Ramage, 2009). Potentiometric-surface declines cause a decrease in hydraulic pressure (depressuring) that creates a load on the skeletal matrix of the sediments in the aquifer and adjacent confining units (fig. 5). Because sand layers are more transmissive and less compressible than clay layers, sand layers depressure more rapidly compared to clay layers. In addition, when groundwater withdrawals are decreased, pressure equilibrium is reestablished more rapidly in the sand layers compared to the clay layers, and the amount of compaction of the sand layers is usually minor compared to the amount of compaction of the clay layers (Trahan, 1982; Galloway and others, 1999). Depressurizing causes a slower dewatering of the clay layers compared to the sand layers; the clay layers often are interbedded with the sand layers. The compressibility of the clay layers is dependent on the thickness and hydraulic characteristics of the clay layers and the vertical stress of the sediment overburden. Slow drainage of the clay layers continues to occur until the excess residual pore pressure in the clay layers equilibrates with the pore pressure of the adjacent sand layers. As dewatering progresses, compaction of the clay layers continues until pressure equilibrium is attained. A similar loading process can occur in sand layers; however, a major difference is that the orientation of the individual clay grains realign as dewatering progresses, becoming perpendicular to the applied vertical overburden load (Galloway and others, 1999). Therefore, the dewatering caused by the



**Figure 5.** Diagram depicting mechanism of subsidence in an aquifer composed of sand and clay (modified from Galloway and others, 1999, p. 9).

depressurization of the clay layers, combined with clay-grain realignment, reduces the porosity and groundwater-storage capacity of the clay layers as compaction occurs (fig. 5). Because most compaction of the clay layers is inelastic, about 90 percent of the compaction is permanent (Gabrysch and Bonnet, 1975). In areas where groundwater withdrawals have decreased, the water level in the aquifers rises and represses the compacted clay layers, but only a small amount of rebound of the land-surface altitude occurs (Gabrysch and Bonnet, 1975). Although the compaction of one thin clay layer generally will not cause a noticeable decrease in the land-surface altitude, when numerous stacked clay- and sand-layer sequences (characteristic of the Gulf Coast aquifer system) depressure and compact, subsidence often occurs (Gabrysch and Bonnet, 1975).

## Data Collection and Analysis Methods

Water-level data were obtained from wells by measuring the depth to water below land surface at each well. Measurements were obtained using calibrated steel tape, air line, and electric water-level tape and also were obtained from reports of well operators. Most wells are pumped at least once daily, but some are pumped more frequently. A minimum of two water-level measurements were made at each well while the well was not being pumped; however, antecedent pumping conditions and pumping status of nearby wells were not always known. To ensure that the water-level measurement recorded for each well is accurate, additional water-level measurements are made as required. Water-level-measurement data used to create sheets 1–14 of the report were collected during December 2009–March 2010 (tables 1–14); during winter months, water levels in the Houston-Galveston region are usually higher compared to the rest of the year because rates of groundwater withdrawals are at a minimum. Subsequently, these data were incorporated into a Geographic Information System (GIS) as point-data layers and used for the construction of sheets 1–14.

### Water-Level Altitudes

The water-level-altitude data used to construct the water-level-altitude maps for the Chicot, Evangeline, and Jasper aquifers (sheets 1, 6, and 11, respectively) were determined by subtracting the water-level measurement from the land-surface-altitude value for each well (point). Land-surface-altitudes were referenced to either the National Geodetic Vertical Datum of 1929 (NGVD 29) or the North American Vertical Datum of 1988 (NAVD 88) (National Oceanic and Atmospheric Administration, 2010). Therefore, each point used for contour configuration on the three water-level-altitude maps is referenced to either NGVD 29 or NAVD 88 (tables 1, 6, 11). Water-level altitudes are depicted with contour intervals of 25, 50, and 100 ft.

### Changes in Water-Level Altitudes

Maps depicting changes in water-level altitudes in the Chicot, Evangeline, and Jasper aquifers were prepared for 2009–10, 2005–10, 1990–2010 (Chicot and Evangeline), 1977–2010 (Chicot and Evangeline), and 2000–10 (Jasper). To create the various water-level-change maps, datasets of water-level-change values (difference between the current year [2010] and historical water-level-altitude values) are used.

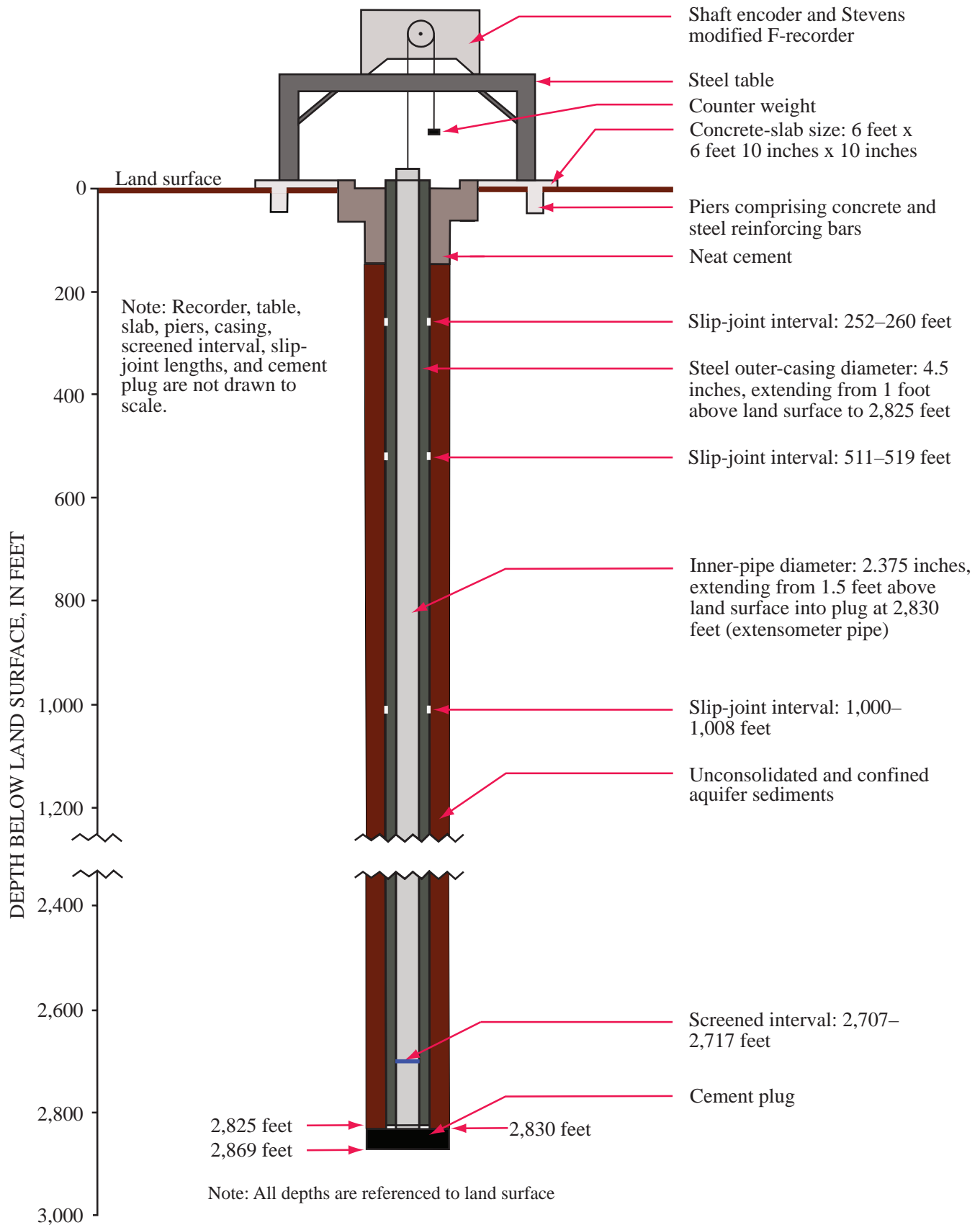
For the 1-year (2009–10) water-level-change maps (sheets 2, 7, and 12), water-level-change data were computed as the difference between water-level altitude at each point (well) for which a water-level measurement was made in 2009 and in 2010. Water-level changes are shown as a blue triangle (water-level rise), a red inverted triangle (water-level decline), or a circle (no water-level change).

For the 5-year (2005–10) change maps (sheets 3, 8, and 13), water-level changes were computed the same as for the 1-year maps—the difference between water-level altitude at each well for which a water-level measurement was made in 2010 and 2005. Changes on the 5-year maps are indicated by contours of equal water-level change. Each 5-year map was prepared by contouring the set of mapped point differences.

For the historical (1990–2010, 1977–2010, 2000–10) change maps (sheets 4, 5, 9, 10, and 14), water-level changes were computed as the difference between water-level altitude at each well for which a water-level measurement was made in 2010 and in the historical year. For wells measured in 2010 that had no corresponding measurement in the historical year, the water-level altitude was estimated using GIS methods for the point represented by that well. A GIS raster (gridded surface) (Worboys, 1995) was created from published historical water-level-altitude contours (1990, 1977, and 2000), and the current-year GIS well locations were intersected with the historical gridded surface to estimate the water-level altitude. Subsequently, the two water-level altitudes for the point were subtracted, creating a water-level-change value (control-point data) that was used to develop water-level-altitude contours. This GIS technique was necessary because numerous wells measured in 1990, 1977, and 2000 either have been destroyed or could not be measured in 2010.

### Borehole Extensometers

To install a borehole extensometer, a borehole is first drilled to a predetermined depth, generally below the depth of expected water-level decline. A steel outer casing with slip joints is then installed in the borehole from 1 ft above land surface to 5 ft above a cement plug placed at the bottom of the borehole (total depth of the borehole is from land surface to cement plug), and a smaller diameter pipe (inner pipe; often referred to as the extensometer pipe) is inserted inside the outer casing, with the inner pipe terminating within the cement plug and extending above land surface (fig. 6). The slip joints



**Figure 6.** Diagram depicting cross-sectional perspective of the borehole extensometer/piezometer (LJ-65-23-322) at Pasadena, Texas.

in the outer casing prevent crumpling and collapse of the outer casing as clay compaction associated with a decrease in land-surface altitude occurs, while the inner pipe remains fixed and rigid. At land surface, a concrete slab is poured and connected to an array of vertical concrete piers extending down to the water table. This construction design helps eliminate the continuous shrinking and swelling of the clayey surficial sediments associated with soil-moisture changes. The concrete piers connect the slab to the underlying unconsolidated sediments penetrated by the borehole. A gage house is constructed on the slab, and a steel-table-mounted recorder is installed in the gage house. A calibrated steel tape connects the recorder to the top of the inner pipe; because the steel table is anchored to the slab, changes in land-surface altitude can be accurately measured and recorded. These recorded values represent the cumulative clay compaction at the site. The scientific theory and operation of a borehole extensometer is further explained by Gabrysch (1984).

Extensometer data are used to quantify the rate of compaction in aquifer formations, providing water-resource managers a tool for evaluating the effects on subsidence rates caused by changes in the amount of groundwater withdrawn from the Chicot and Evangeline aquifers. For this report, borehole extensometer data measuring the compaction of subsurface material in the Chicot and Evangeline aquifers were evaluated at 13 borehole extensometers in Harris and Galveston Counties. To quantify the rate of compaction in aquifer formations at selected sites throughout Harris and Galveston Counties, a network of extensometers was installed beginning in 1973. Five extensometers were installed in Harris or Galveston County in July 1973: KH-64-33-920 (Texas City-Moses Lake) in Galveston County; LJ-65-22-622 (East End), LJ-65-16-930 (Baytown C-1), LJ-65-16-931 (Baytown C-2), and LJ-65-32-625 (Seabrook) in Harris County. A previously existing borehole extensometer installed in 1962 in Harris County (LH-65-32-401 [Johnson Space Center]) was included in the network, and routine measurements of compaction obtained at the Johnson Space Center since July 1973 are included in this report. Additional extensometers were added to the network during 1974–76 in Harris County: LJ-65-12-726 (Addicks) in 1974, LJ-65-23-322 (Pasadena) in 1975, and LJ-65-32-428 and LJ-65-32-424 (Clear Lake) in 1976. The last three extensometers in the current (2010) network were installed in Harris County in 1980: LJ-65-07-909 (Lake Houston), LJ-65-14-746 (Northeast), and LJ-65-21-226 (Southwest). Since activation or installation between 1973 and 1980, compaction measurements have been obtained at these 13 extensometers at least monthly, providing site-specific rates of compaction accurate within 0.001 ft.

Each borehole extensometer has a 10-ft open interval at depth and above the cement plug allowing water to flow into the center pipe and thus functions as a piezometer (small-diameter well used to measure water-level altitudes). A water-level measurement is made during each extensometer site visit. If the depth of the screened interval is in the Chicot or

Evangeline aquifer, these water-level measurements are considered during construction of the annual water-level-altitude map for the Chicot and Evangeline aquifers, respectively.

## Quality Assurance

Protocols for the collection and review of water-level-altitude data were followed as described in the USGS Texas Water Science Center “Quality Assurance Plan for Groundwater Activities” (appendix 7.3, “Groundwater Data Management Plan” [G.P. Stanton, U.S. Geological Survey, written commun., 2010]). All data were stored in the USGS National Water Information System (NWIS) (U.S. Geological Survey, 2010).

The annual (2010) regional depictions of water-level altitudes presented in this report are derived from water-level-measurement data collected during December 2009–March 2010 throughout an 11-county area that includes the greater Houston-Galveston area. The water-level altitudes of the Chicot, Evangeline, and Jasper aquifers change in response to changes in hydrologic conditions and groundwater pumping rates; water-level-altitude changes might have occurred since the most recent water-level measurements were made. Antecedent withdrawal rates and pumping status of nearby wells were not always known and could have affected the representativeness of the water-level data that were collected.

## Water-Level Altitudes and Changes

Locations of wells used to construct the water-level-altitude and water-level-change maps for the Chicot, Evangeline, and Jasper aquifers are shown in appendix 1. The well index numbers on the three maps (appendixes 1.1, 1.2, 1.3) coincide with tabular data (tables 1–14) for each water-level altitude or water-level-change map, and the well symbols (small circles) show the geographic location of the well.

The Chicot and Evangeline aquifer maps show approximate water-level altitudes in 2010 and water-level changes for 2009–10, 2005–10, 1990–2010, and 1977–2010 (sheets 1–5 and 6–10, respectively). Depictions of long-term water-level change (1977–2010) show areas of decline in northern, northwestern, and southwestern Harris County and a broad area of water-level rise in southeastern Harris and northern Galveston Counties.

The Jasper aquifer maps show approximate water-level altitudes in 2010 and water-level changes for 2009–10, 2005–10, and 2000–10 (sheets 11–14). Depictions of long-term water-level change (2000–10) show that water-level declines encompass most of Montgomery County and north-central Harris County.

Except for the 2009–10 change maps, the water-level-altitude contours were prepared using contour intervals specific to the range of water-level-altitude changes for a



given map. Adjusting the contour intervals in this way helped present a clear depiction of regional water-level-altitude changes.

## **Chicot Aquifer**

Water-level measurements from 180 wells (table 1) were used to construct the 2010 water-level-altitude map of the Chicot aquifer. In 2010, water-level-altitude contours for the Chicot aquifer ranged from 200 ft below NGVD 29 or NAVD 88 (hereinafter, datum) in a small area in southwestern Harris County to 200 ft above datum in central to southwestern Montgomery County (sheet 1). The Chicot aquifer water-level-change maps for 2009–10, 2005–10, 1990–2010, and 1977–2010 are shown on sheets 2, 3, 4, and 5, respectively. The number of water-level measurement pairs used to construct the change maps were 171 for 2009–10 (table 2), 144 for 2005–10 (table 3), 148 for 1990–2010 (table 4), and 144 for 1977–2010 (table 5).

Changes in water-level altitudes in the Chicot aquifer for 2009–10 ranged from a 49-ft decline in southwestern Montgomery County to a 67-ft rise in southwestern Harris County (sheet 2) and for 2005–10 ranged from a 25-ft decline in western Harris County to a 35-ft rise in areas of southeastern Harris County and the northern boundary between Fort Bend and Brazoria Counties (sheet 3). Changes in water-level altitudes for 1990–2010 ranged from a 40-ft decline in east-central Fort Bend and south-central Harris Counties to an 80-ft rise in eastern Harris County (sheet 4) except for an isolated area in south-central Harris County with a 180-ft rise that represents a few points; however, because of the close proximity of these points and the large differences between these and other nearby water-level-altitude changes, these few points were not used to construct contours representing regional water-level change in the area. For 1977–2010, changes in water-level altitudes ranged from a 140-ft decline in north-central Harris County to a 200-ft rise in southeastern Harris County (sheet 5).

## **Evangeline Aquifer**

Water-level measurements from 356 wells (table 6) were used to construct the 2010 water-level-altitude map of the Evangeline aquifer. In 2010, water-level-altitude contours for the Evangeline aquifer ranged from 300 ft below datum in north-central Harris County to 200 ft above datum at the boundary of Waller, Montgomery, and Grimes Counties (sheet 6). The Evangeline aquifer water-level-change maps for 2009–10, 2005–10, 1990–2010, and 1977–2010 are shown on sheets 7, 8, 9, and 10, respectively. The number of water-level measurement pairs used to construct the change maps were 330 for 2009–10 (table 7), 275 for 2005–10 (table 8), 294 for 1990–2010 (table 9), and 291 for 1977–2010 (table 10).

Changes in water-level altitudes in the Evangeline aquifer for 2009–10 ranged from a 58-ft decline in north-central

Harris County to a 69-ft rise in southwestern Harris County (sheet 7) and for 2005–10 ranged from an 80-ft decline in northwestern Harris County to an 80-ft rise in southwestern Harris County (sheet 8). Changes in water-level attitudes for 1990–2010 ranged from a 200-ft decline in south-central Montgomery County to a 220-ft rise in southeastern and central Harris County (sheet 9) and for 1977–2010 ranged from a 320-ft decline at the boundary between Harris and Montgomery Counties to a 220-ft rise in southeastern Harris County (sheet 10).

## **Jasper Aquifer**

Water-level measurements from 93 wells (table 11) were used to construct the 2010 water-level-altitude map of the Jasper aquifer. In 2010, water-level-altitude contours for the Jasper aquifer ranged from 200 ft below datum in south-central Montgomery County to 250 ft above datum in east-central Grimes County (sheet 11).

The Jasper aquifer water-level-change maps for 2009–10, 2005–10, and 2000–10 are shown on sheets 12, 13, and 14, respectively. The number of water-level measurement pairs used to construct the change maps were 78 for 2009–10 (table 12), 55 for 2005–10 (table 13), and 79 for 2000–10 (table 14).

Changes in water-level altitudes in the Jasper aquifer for 2009–10 ranged from a 39-ft decline in north-central Montgomery County to a 39-ft rise in San Jacinto County (sheet 12) and for 2005–10 ranged from a 110-ft decline in northwestern-central Harris County to no change in far eastern Montgomery and north-central Harris Counties (sheet 13). For 2000–10, changes in water-level altitudes ranged from a 180-ft decline in south-central Montgomery County to no change in far northwestern Montgomery County (sheet 14).

## **Compaction in the Chicot and Evangeline Aquifers**

Compaction of subsurface materials (mostly of the clay layers because little compaction occurs in sand layers; refer to “Subsidence and Compaction Processes” section) composing the Chicot and Evangeline aquifers was recorded continuously at the 13 borehole extensometers at 11 sites (sheet 15) using analog chart recorders. The rate of compaction varies from site to site (sheet 16). Graphs of compaction are shown for 1973, or later, through 2009 (depending on when each extensometer was installed) for 12 of the 13 extensometers (sheet 16); compaction data for the graphs on sheet 16 are listed in tables 15A–L. Compaction measured by the proximal and shallower of the two borehole extensometers at the Clear Lake site (LJ-65-32-424) is not shown because the recorded data are similar to that measured by the deeper borehole extensometer (LJ-65-32-428) at the site. The depth of the extensometer determines the total thickness of sediment where clay compaction is measured by the extensometer. Five extensometers

measure clay compaction in the sediments of the Chicot aquifer (East End [LJ-65-22-622], Johnson Space Center [LH-65-32-401], Texas City-Moses Lake [KH-64-33-920], Baytown C-1 [LJ-65-16-930], Seabrook [LJ-65-32-625]), and seven extensometers measure compaction in the sediments of the Chicot and Evangeline aquifers (Lake Houston [LJ-65-07-909], Northeast [LJ-65-14-746], Southwest [LJ-65-21-226], Addicks [LJ-65-12-726], Baytown C-2 [LJ-65-16-931], Clear Lake [LJ-65-32-428], Pasadena [LJ-65-23-322]) (sheet 16). Since the early 1900s, as much as 12–13 ft of historical subsidence has occurred in the Pasadena and Baytown areas in Harris County (Kasmarek, Gabrysch, and Johnson, 2009). Most of this subsidence occurred before the 1970s; the graphs of cumulative-compaction data from installation in 1975 through 2009 for the Pasadena extensometer and from installation in 1973 through 2009 for the Baytown extensometers indicate subsidence of 0.250, 0.650, and 0.954 ft, respectively (sheet 16; tables 15L, 15H, and 15I, respectively). Most of the subsidence (77 to 97 percent) in the Houston-Galveston region occurred prior to 1973 when the installation of extensometers began (Kasmarek, Gabrysch, and Johnson, 2009).

Prior to the creation of HGSD, the withdrawal of groundwater from the Chicot and Evangeline aquifers was unregulated, and the associated compaction of clay layers was ongoing as water levels in the aquifers were declining. By 1977, the withdrawals had resulted in water-level-altitude declines of 300 and 350 ft below datum in the Chicot and Evangeline aquifers, respectively, in southeastern Harris County (Gabrysch, 1979). The rate of compaction varies from site to site because of differences in groundwater withdrawals near each site and differences among sites in the clay-to-sand ratio of the subsurface materials. When reductions in groundwater withdrawals were first mandated following the creation of the HGSD in 1975, the rate of groundwater withdrawal began to decrease along with the rate of clay compaction (sheet 16). Coincident with the curtailment of groundwater withdrawals, the aquifer water levels began to rise. Water levels in the Chicot and Evangeline aquifers have risen as much as 200 and 220 ft, respectively, as shown on the 1977–2010 long-term-change maps (sheets 5 and 10, respectively).

For the period of record ending December 2009, cumulative clay compaction in the Chicot and Evangeline aquifers (sheet 16) ranged from 0.088 ft (table 15G) at the Texas City-Moses Lake extensometer to 3.559 ft (table 15E) at the Addicks extensometer. The graphs of cumulative-compaction data show that the compaction rates were appreciably higher when the extensometers were initially installed compared to compaction rates in subsequent years. These asymptotic compaction-rate decreases were directly related to the rise in water levels in the aquifers as groundwater withdrawals were decreased in response to regulatory mandates of the HGSD. As the water level of the aquifers began to rise or rebound, the hydrostatic pressure increased and excess, residual pore

pressure equilibrated, hence the rates of compaction progressively decreased. Coinciding with compaction-rate decreases, the long-term water-level changes for 1990–2010 and 1977–2010 in the Chicot and Evangeline aquifers (sheets 4 and 5 and sheets 9 and 10, respectively) show that, except for the Addicks extensometer, the locations of these extensometers coincide with the relatively large area of water-level rise. Compaction data from the Addicks extensometer (table 15E) show that the rate of compaction continued steadily through mid-2003. The reason the rate remained steady during this period is that the extensometer is located in Regulatory Area 3 of HGSD, and as such, was not scheduled for a 30-percent groundwater reduction until 2010 (Harris-Galveston Subsidence District, 2010). Therefore, from mid-1974 through mid-2003, groundwater withdrawal continued in the area adjacent to the Addicks extensometer site with an associated clay-compaction rate of about 0.1 ft per year. Additionally, the rate of compaction during about August 2003–December 2003 decreased to about 0.004 ft because an adjacent public-supply well field was inoperative during this period. From December 2003 until March 2005, recorded data indicate an increase (rebound) in land-surface altitude of about 0.030 ft. The graph of compaction for the Seabrook extensometer shows a seasonal sinusoidal trend that indicates the altitude of the land surface decreases in the hot and dry summer months as the surficial clayey sediments desiccate. As the heat of the summer dissipates and the cooler and wetter months arrive, the sediments rehydrate causing the altitude of the land surface to increase and heave upward.

Compaction data for the Texas City-Moses Lake site indicate that not only has the rate of compaction been halted but also, since about 1981, a slight land-surface rise of approximately 0.1 ft has occurred. The graphs of compaction data for the Pasadena, Clear Lake, Seabrook, Baytown C-1 and C-2, and Johnson Space Center extensometers indicate a slight increase in land-surface altitude from late-1978 to early-1980 because of a ruptured natural gas well that pressurized the aquifer system, causing water levels to rise in the area adjacent to the well (Gabrysch, 1984). Over a period of about 2 years, the pressure in the aquifer slowly dissipated and the rates of compaction returned to pre-pressuring-event rates. The graphs of compaction data for the two Baytown extensometers show a noticeable amount of seasonal change from late-1973 to late-1992, most likely caused by the expansive characteristic of the montmorillonitic clay composing the aquifer sediments. When soil moisture is comparatively low in the summer, the surficial clays of the sediments compress, or shrink, causing a decrease in land-surface altitude. Conversely, when soil moisture is comparatively high in the winter, the clays expand, or swell, causing an increase in land-surface altitude. To address the problem of shrinking and swelling clays, a modification was made in 1982 to the original design of the borehole extensometers by installing a system of vertical piers connected to the concrete bases of the extensometers and extending downward to the depth of the water table.

## Data Limitations

Most land-surface altitudes for wells in this report are estimates from USGS 7.5-minute topographic quadrangle maps. Land-surface altitudes for 2009 and later for the wells in Harris County are derived from a Light Detection and Ranging (LiDAR)-based digital elevation model (DEM) (Tropical Storm Allison Recovery Project, 2009). These altitudes are referenced to NGVD 29 and NAVD 88, respectively. The variability between the vertical datums over the 11,000-mi<sup>2</sup> geographic area described in this report ranged from -0.073 to 0.095 ft with a mean difference of -0.003 ft when examined using Corpscon version 6 and Geoid03 (U.S. Army Corps of Engineers, 2005). LiDAR data were contoured at a 1-ft interval, providing 0.5-ft accuracy. The 7.5-minute topographic maps for the Gulf Coast area were normally contoured at a 5-ft interval, thereby providing a 2.5-ft accuracy; thus, the LiDAR data provided about 5 times better accuracy when compared to topographic maps (Kasmarek, Gabrysch, and Johnson, 2009). In addition, the 7.5-minute topographic maps have not been updated with changes in land-surface altitude that might have occurred since publication of the map. The effects of land-surface-altitude changes on water-level-change maps are unavoidable if the change maps are to accurately reflect differences between current-year and previous-year water-level-altitude maps (each of which reflects the best available land-surface altitudes of wells).

The depictions of water-level altitudes and changes at any specific location are considered to represent a regional-scale approximation and, as such, are not intended for use in engineering or other design applications. The water-level measurements collected for this report were rounded to the nearest foot; the values shown on the maps represent a mathematical approximation that could vary as much as  $\pm 0.5$  ft, in addition to accuracies associated with the source data. Use of these data for critical or local-scale applications is not advised without full awareness of the data limitations. Users need to exercise discretion when drawing conclusions or making policy decisions on the basis of these contoured depictions.

Compaction data recorded at each extensometer site are the cumulative compaction for all subsurface materials above the depth of the cement plug (fig. 6); any compaction or vertical movement that occurs below these depths (sheet 16) are not measured by the extensometer. The compaction of subsurface material for a given extensometer could be occurring solely in the sediments of the Chicot aquifer (for example, the Baytown C-1 extensometer) or a combination of the sediments of the Chicot and Evangeline aquifers (for example, the Addicks extensometer). In addition to differences in groundwater withdrawals near each extensometer, the clay-to-sand ratio is different at each site, hence, the rate of compaction varies from site to site (sheet 16). Therefore, it is not possible to extrapolate or infer a rate of clay compaction for an area on the basis of the rate of compaction measured at a nearby extensometer.

## Summary

Groundwater withdrawn from the Chicot, Evangeline, and Jasper aquifers has been the primary source of water for municipal supply and for commercial and industrial uses in the Houston-Galveston region, Texas, since the early 1900s. Most of the subsidence in the Houston-Galveston region has occurred as a direct result of groundwater withdrawals for municipal supply, industrial use, and irrigation that depressed and dewatered the Chicot and Evangeline aquifers causing compaction of the clay layers of the aquifer sediments. By 1979, as much as 10 ft of subsidence had occurred in the Houston-Galveston region, and approximately 3,200 mi<sup>2</sup> of the 11,000-mi<sup>2</sup> geographic area had subsided more than 1 ft. To address the issue of subsidence and subsequent increased flooding, the Texas State Legislature in 1975 authorized the establishment of the Harris-Galveston Subsidence District to regulate and reduce groundwater withdrawals in Harris and Galveston Counties. The Fort Bend Subsidence District subsequently was established in 1989 to regulate groundwater withdrawals in Fort Bend County and the Lone Star Groundwater Conservation District in 2001 to regulate groundwater withdrawals in Montgomery County. This report, prepared by the U.S. Geological Survey in cooperation with the Harris-Galveston Subsidence District, City of Houston, Fort Bend Subsidence District, and Lone Star Groundwater Conservation District, is one in an annual series of reports depicting water-level altitudes and water-level changes in the Chicot, Evangeline, and Jasper aquifers and compaction in the Chicot and Evangeline aquifers in the Houston-Galveston region.

The report contains maps showing 2010 water-level altitudes for the Chicot, Evangeline, and Jasper aquifers, respectively; maps showing 1-year (2009–10) water-level-altitude changes for each aquifer; maps showing 5-year (2005–10) water-level-altitude changes for each aquifer; maps showing long-term (1990–2010 and 1977–2010) water-level-altitude changes for the Chicot and Evangeline aquifers; a map showing long-term (2000–10) water-level-altitude change for the Jasper aquifer; a map showing locations of borehole extensometer sites; and graphs showing measured compaction of subsurface material at the extensometers from 1973, or later, through 2009. Tables listing the data used to construct each aquifer-data map and the compaction graphs are included.

Water levels in the Chicot, Evangeline, and Jasper aquifers were measured during December 2009–March 2010 when water levels usually are higher compared to the rest of the year. Water-level measurements from 180 wells were used to construct the 2010 water-level-altitude map of the Chicot aquifer; contours ranged from 200 ft below NGVD 29 or NAVD 88 (hereinafter, datum) in a small area in southwestern Harris County to 200 ft above datum in central to southwestern Montgomery County. Water-level-altitude changes in the Chicot aquifer for 2009–10 ranged from a 49-ft decline



to a 67-ft rise; for 2005–10 ranged from a 25-ft decline to a 35-ft rise; for 1990–2010 ranged from a 40-ft decline to an 80-ft rise; and for 1977–2010 ranged from a 140-ft decline to a 200-ft rise. Water-level measurements from 356 wells were used to construct the 2010 water-level-altitude contours of the Evangeline aquifer; contours ranged from 300 ft below datum in north-central Harris County to 200 ft above datum at the boundary of Waller, Montgomery, and Grimes Counties. Water-level-altitude changes in the Evangeline aquifer for 2009–10 ranged from a 58-ft decline to a 69-ft rise, for 2005–10 ranged from an 80-ft decline to an 80-ft rise, for 1990–2010 ranged from a 200-ft decline to a 220-ft rise, and for 1977–2010 ranged from a 320-ft decline to a 220-ft rise. Water-level measurements from 93 wells were used to construct the 2010 water-level-altitude contours of the Jasper aquifer; contours ranged from 200 ft below datum in south-central Montgomery County to 250 ft above datum in eastern-central Grimes County. Water-level-altitude changes in the Jasper aquifer for 2009–10 ranged from a 39-ft decline to a 39-ft rise, for 2005–10 ranged from a 110-ft decline to no change, and for 2000–10 ranged from a 180-ft decline to no change. For the Chicot and Evangeline aquifers, depictions of long-term water-level change (1977–2010) show areas of decline in northern, northwestern, and southwestern Harris County and a broad area of water-level rise in southeastern Harris and northern Galveston Counties. For the Jasper aquifer, depictions of long-term water-level change (2000–10) show that water-level declines encompass most of Montgomery County and north-central Harris County.

Compaction of subsurface materials (mostly in the clay layers) composing the Chicot and Evangeline aquifers was recorded continuously at the 13 borehole extensometers at 11 sites. The rates of compaction measured by each extensometer were generally largest in the first few years after installation. When reductions in groundwater withdrawals were first required following the creation of the Harris-Galveston Subsidence District in 1975, the rate of groundwater withdrawal began to decrease along with the rate of clay compaction. Coincident with the curtailment of groundwater withdrawals, the water levels of the aquifers began to rise. Water levels in the Chicot and Evangeline aquifers have risen as much as 200 and 220 ft, respectively, as shown on the 1977–2010 long-term-change maps. For the period of record beginning in 1973, or later, and ending in December 2009, cumulative clay compaction at 12 of the 13 borehole extensometers in the Chicot and Evangeline aquifers ranged from 0.088 ft at the Texas City-Moses Lake extensometer (KH-64-33-920) to 3.559 ft at the Addicks extensometer (LJ-65-12-726). The rate of compaction varies from site to site because of differences in groundwater withdrawals near each site and differences among sites in the clay-to-sand ratio of the subsurface materials. Therefore, it is not possible to extrapolate or infer a rate of clay compaction for an area on the basis of the rate of compaction measured at a nearby extensometer.

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