



**DISCUSSION**

This sheet includes maps that show the thickness and the depth to base of uppermost Pleistocene and Holocene (in other words, post-Late Glacial Maximum) deposits for the Offshore of Carpinteria map area (Maps A, B), as well as for a larger area that extends about 115 km along the coast from Huene Canyon to Refugio Beach (Maps D, E) to establish a regional context. To make these maps, water bottom and depth to base of the post-Late Glacial Maximum horizons were mapped from seismic-reflection profiles. The difference in the two horizons was reported for every point on an XY coordinates (UTM zone 11) and two-way travel time (TWT). The thickness of the post-Late Glacial Maximum unit (Maps B, E) was determined by applying a sound velocity of 1400 m/s to the TWT. The thickness points were interpolated to a preliminary continuous surface, overlaid with zero-thickness bedrock outcrops (sheet 10 of this report, and contoured (Wong and others, 2012). Data within Huene Canyon were excluded from the contouring because the seismic-reflection data are too sparse to adequately image the highly variable changes in sediment thickness that characterize the canyon (Maps A, B, D, E).

Several factors required manual editing of the preliminary thickness maps to make the final products. The Red Mountain Fault Zone (RMFZ), Pitas Point Fault (PPF), and Oak Ridge Fault (ORF) distort the sediment sequence in the region (Maps D, E). The data points are dense along tracklines (about 1 m apart) and sparse between tracklines (1-2 km apart), resulting in contoured artifacts. To incorporate the effect of the faults, to remove irregularities from interpolation, and to reflect other geologic information and complexity, the resulting interpolated contours were modified. Contour modifications and regrading were repeated several times to produce the final regional sediment-thickness map (Wong and others, 2012).

Data to determine the depth to base of the post-Late Glacial Maximum unit was similarly processed and contoured. However, this preliminary data set was set aside in favor of a surface determined by subtracting the modified thickness data from multibeam bathymetry collected separately (see sheet 11). The depth of this surface in the Huene Canyon to Refugio Beach area ranges from 12 to 190 m (Map D; see also, Wong and others, 2012).

Five "arenas" of sediment thickness, which are bounded either by faults or by Huene Canyon, are recognized on the regional maps (Maps D, E): (1) Refugio Beach to the south strand of the Red Mountain Fault Zone (RMFZ), (2) between the south strand of the Red Mountain Fault Zone and the Pitas Point Fault (PPF), (3) between the Pitas Point Fault and Oak Ridge Fault (ORF), (4) between the Oak Ridge Fault and Huene Canyon, and (5) south of Huene Canyon. These data highlight the contrast among three general zones of sediment thickness: (1) the uplifted, sediment-poor Santa Barbara shelf domain 1; mean sediment thickness of 3.5 m; (2) a transitional zone (domain 2); mean sediment thickness of 18 m; and (3) the subsiding, sediment-rich delta and shelf offshore of the Ventura and Santa Clara Rivers and Calleguas Creek (domains 3, 4, and 5); mean sediment thicknesses of 39.3, 38.9, and 28.3 m, respectively.

In the Offshore of Carpinteria map area, sediment thickness ranges from 0 to 47 m, with a mean thickness of 12.7 m. The thickest sediment accumulation (mean thickness, 44.0 m) is south of the north-dipping Pitas Point Fault. Sediment thus markedly (from 40 to 5 m; mean thickness, 19.2 m) from south to north in the uplifted and tilted zone between the Pitas Point Fault and the south strand of the Red Mountain Fault Zone. Sediment is much thinner north of the south strand of the Red Mountain Fault Zone (mean thickness, 5.1 m) where the thickest accumulations are present in a gentle trough between the uplifts associated with the Red Mountain and Rincon Creek Faults (see fig. 1 on sheet 8). The origin of this trough could be either structural or erosional (or both); although it is spatially related to a broad syncline (see figs. 1, 2 on sheet 8), it also could represent the axis of a lowstand paleodrainage system.

The regional pattern of faults and of earthquakes occurring between 1922 and 2010 that have informed or measured magnitudes greater than 2.0 is shown on Map C. Fault locations are based on our mapping within California's State Waters and on a generalized compilation of mapping from Heck (1998), Moore and others (2009), and Jennings and Bryant (2010). Earthquake data are from the Southern California Earthquake Catalog (Southern California Earthquake Data Center, 2010). Although earthquake locations have been determined by the Caltech network since 1922, significantly greater precision began in 1969 with installation of a U.S. Geological Survey (USGS) seismicographic network (Lee and Vidale, 1973; Southern California Earthquake Data Center, 2010). Epicentral data indicate that recent seismicity in the eastern and central Santa Barbara Channel is characterized by earthquake swarms, relatively frequent minor earthquakes, and infrequent major earthquakes.

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**Local (Offshore of Carpinteria Map Area) and Regional (Offshore from Refugio Beach to Huene Canyon) Shallow-Subsurface Geology and Structure, Santa Barbara Channel, California**

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