

## Geologic Map of the Granite 7.5' Quadrangle, Lake and Chaffee Counties, Colorado



Pamphlet to accompany  
Scientific Investigations Map 3294

**U.S. Department of the Interior**  
**U.S. Geological Survey**

**COVER.** Oblique aerial view to the northwest looking toward the Twin Lakes Reservoir and the snow-capped peaks in the Sawatch Range. Highest peak on the skyline [far right] is Mount Elbert, 4,399 m (14,433 ft). Tree- and sagebrush-covered ridges on the south and north sides of the Twin Lakes reservoir are moraines composed of till of Pinedale, Bull Lake, and pre-Bull Lake age. Arkansas River and US Highway 24 are in foreground. Photograph by Jason P. Briner, University at Buffalo, Buffalo, New York, June 8, 2010.

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By Ralph R. Shroba, Karl S. Kellogg, and Theodore R. Brandt

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**U.S. Department of the Interior**  
SALLY JEWELL, Secretary

**U.S. Geological Survey**  
Suzette Kimball, Acting Director

U.S. Geological Survey, Reston, Virginia: 2014

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## Conversion Factors

SI to Inch/Pound

<b>Multiply</b>	<b>By</b>	<b>To obtain</b>
	Length	
centimeter (cm)	0.3937	inch (in.)
millimeter (mm)	0.03937	inch (in.)
meter (m)	3.281	foot (ft)
kilometer (km)	0.6214	mile (mi)
	Area	
square kilometer (km <sup>2</sup> )	0.3861	square mile (mi <sup>2</sup> )
	Flow rate	
cubic meter per second (m <sup>3</sup> /s)	35.31	cubic foot per second (ft <sup>3</sup> /s)

Vertical coordinate information is referenced to the North American Vertical Datum of 1929 (NAVD 29)

Horizontal coordinate information is referenced to the North American Datum of 1927 (NAD 27)

Altitude, as used in this report, refers to distance above the vertical datum.

## Division of Quaternary, Neogene, and Paleogene Time used in this report<sup>1</sup>

<b>Period or subperiod</b>	<b>Epoch</b>		<b>Age</b>
	Holocene		0–11.7 ka
Quaternary	Pleistocene	late	11.7–132 ka
		middle	132–788 ka
		early	788 ka–2.588 Ma
Neogene	Pliocene		2.588–5.332 Ma
	Miocene		5.332–23.03 Ma
	Oligocene		23.03–33.9
Paleogene	Eocene		33.9–55.8
	Paleocene		55.8–65.5 Ma

<sup>1</sup>Ages of time boundaries are those of the U.S. Geological Survey Geologic Names Committee (2010) except those for the late-middle Pleistocene boundary and middle-early Pleistocene boundary, which are those of Richmond and Fullerton (1986). Ages are expressed in ka for kilo-annum (thousand years) and Ma for mega-annum (million years).

## North American Land Mammal Ages used in this report<sup>1</sup>

<b>Land Mammal Age</b>	<b>Age</b>
Clarendonian	8.9–11.6 Ma
Barstovian	11.6–15.9 Ma
Hemingfordian	15.9–19 Ma
Arikareean	19–29 Ma

<sup>1</sup>Ages of time boundaries are those of Woodburne and Swisher (1995).



# Geologic Map of the Granite 7.5' Quadrangle, Lake and Chaffee Counties, Colorado

By Ralph R. Shroba, Karl S. Kellogg, and Theodore R. Brandt

## Abstract

The geologic map of the Granite 7.5' quadrangle, Lake and Chaffee Counties, Colorado, portrays the geology in the upper Arkansas valley and along the lower flanks of the Sawatch Range and Mosquito Range near the town of Granite. The oldest rocks, exposed in the southern and eastern parts of the quadrangle, include gneiss and plutonic rocks of Paleoproterozoic age. These rocks are intruded by younger plutonic rocks of Mesoproterozoic age. Felsic hypabyssal dikes, plugs, and plutons, ranging in age from Late Cretaceous or Paleocene to late Oligocene, locally intruded Proterozoic rocks. A small andesite lava flow of upper Oligocene age overlies Paleoproterozoic rock, just south of the Twin Lakes Reservoir. Gravelly fluvial and fan deposits of the Miocene and lower Pliocene(?) Dry Union Formation are preserved in the post-30 Ma upper Arkansas valley graben, a northern extension of the Rio Grande rift. Mostly north-northwest-trending faults displace deposits of the Dry Union Formation and older rock units. Light detection and ranging (lidar) imagery suggests that two short faults, near the Arkansas River, may displace surficial deposits as young as middle Pleistocene.

Surficial deposits of middle Pleistocene to Holocene age are widespread in the Granite quadrangle, particularly in the major valleys and on slopes underlain by the Dry Union Formation. The main deposits are glacial outwash and post-glacial alluvium; mass-movement deposits transported by creep, debris flow, landsliding, and rockfall; till deposited during the Pinedale, Bull Lake, and pre-Bull Lake glaciations; rock-glacier deposits; and placer-tailings deposits formed by hydraulic mining and other mining methods used to concentrate native gold.

Hydrologic and geologic processes locally affect use of the land and locally may be of concern regarding the stability of buildings and infrastructure, chiefly in low-lying areas along and near stream channels and locally in areas of moderate to steep slopes. Low-lying areas along major and minor streams are subject to periodic stream flooding. Mass-movement deposits and deposits of the Dry Union Formation that underlie moderate to steep slopes are locally subject to

creep, debris-flow deposition, and landsliding. Proterozoic rocks that underlie steep slopes are locally subject to rockfall.

Sand and gravel resources for construction and other uses in and near the Granite quadrangle are present in outwash-terrace deposits of middle and late Pleistocene age along the Arkansas River and along tributary streams in glaciated valleys.

## Introduction

Surficial deposits within the Granite 7.5' quadrangle record alluvial, glacial, mass-movement, and lacustrine processes along the eastern flank of the Sawatch Range, upper Arkansas River valley, and the west flank of the Mosquito Range during the Quaternary. The quadrangle contains one of the most complete sequences of glacial till and outwash in central Colorado. These deposits and their associated landforms have been the subject of reconnaissance studies and detailed investigations since the early 1870s (Hayden, 1874; Capps and Leffingwell, 1904; Davis, 1905; Westgate, 1905; Capps, 1909; Ray, 1940; Richmond and Tweto, 1965; Tweto and Case, 1972; Shroba, 1977; Nelson and Shroba, 1998).

Many of the surficial deposits and deposits of the Dry Union Formation (Nd) are poorly exposed within the map area. Deposits that are limited in extent were not mapped, including fill material along most roads and at most building sites, small deposits of placer tailings, and thin discontinuous colluvial deposits that commonly mantle the Paleogene and older bedrock units. Fractional map symbols (Qsw/Nd and Qc/Nd) are used to show where thin deposits of sheetwash alluvium or colluvium mantle sediments of the Dry Union Formation (Nd). These fractional units are defined in the descriptions of individual units that compose the fractional units, and are not shown on the correlation of map unit diagram.

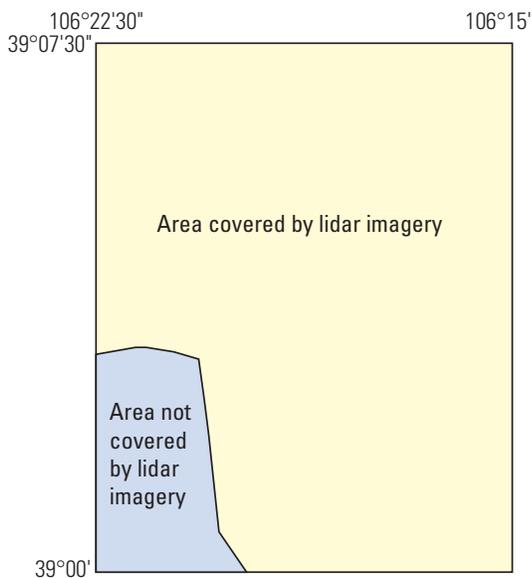
The surficial units on this map are informal allostratigraphic units (disconformity-bound sequences) of the North American Stratigraphic Code (North American Commission on Stratigraphic Nomenclature, 2005), whereas the other map units (bedrock units) are formal lithostratigraphic or informal lithodemic units.

## 2 Geologic Map of the Granite 7.5' Quadrangle, Lake and Chaffee Counties, Colorado

Ages for surficial deposits within the map area are based chiefly on the following: relative heights of alluvial deposits above present stream channels, spatial relationships among glacial till and outwash deposits, relative degree of erosional modification of depositional landforms, degree of soil development on landforms, and clast weathering on and in deposits (see Nelson and Shroba, 1998).

The mapped distribution of surficial map units and the Dry Union Formation is based on field mapping and on interpretation of: (1) light detection and ranging (lidar) imagery generated from lidar data acquired during September 2010 that covers about 88 percent of the Granite 7.5' quadrangle (see fig. 1); (2) 10-foot contour map generated from the lidar data; (3) National Agriculture Imagery Program (NAIP), <http://www.fsa.usda.gov/FSA/apfoapp?area=home&subject=prog&topic=nai> aerial imagery taken in 2011 for about 12 percent the of the Granite quadrangle not covered by lidar data (fig. 1), (4) 1:28,000-scale, NAIP natural color aerial photographs taken in September 2004 (<http://www.fsa.usda.gov>); and (5) 1:40,000-scale, black-and-white aerial photographs taken in September and October 1999 (National Aerial Photography Program, NAPP, <http://www.usgs.gov/science/cite-view.php?cite=553>).

Typical particle sizes of surficial deposits and sediments of the Dry Union Formation in map-unit descriptions are



**Figure 1.** Area of Granite 7.5' quadrangle covered by light detection and ranging (lidar) imagery shown in yellow. This imagery, along with field observations, was used to map surficial deposits, Dry Union Formation, inferred faults, lineaments, fluvial scarps, and other geomorphic features. Area in blue indicates region of quadrangle not covered by lidar imagery data that was interpreted with the aid of National Agriculture Imagery Program (NAIP) imagery.

field estimates. Size limits for sand (2–0.05 mm), silt (0.05–0.002 mm), and clay (<0.002 mm) in table 1 are those of the Soil Survey Staff (1951). In descriptions of surficial deposits and the Dry Union Formation, the term “clasts” refers to granules and larger particles (>2 mm in diameter), whereas the term “matrix” refers to sand and finer particles (≤2 mm in diameter). In descriptions of clast composition of surficial deposits and the Dry Union Formation, the terms “granitic” refer chiefly to rocks ranging in composition from granite to granodiorite, the term “gneiss” refers chiefly to biotite gneiss and hornblende gneiss. In this report, the terms “stream alluvium” and “fluvial deposits” refer to sediment transported by running water confined to channels, whereas the terms “sheetwash alluvium” and “sheetflood alluvium” refer to sediment transported by running water not confined to channels. The terms “colluvium” and “mass-movement deposits” refers to all rock and sediment transported downslope chiefly by gravity (Hilgard, 1892; Merrill, 1897). Material on slopes is transported chiefly by mass-movement (gravity-driven) processes—such as creep, debris flow, sliding, and rockfall—that are locally aided by running water not confined to channels.

Map units that include debris-flow deposits, such as some of the surficial map units and the Dry Union Formation, probably also include hyperconcentrated-flow deposits. These latter deposits have bedding and sorting characteristics intermediate between those of stream-flow and debris-flow deposits (Pierson and Costa, 1987; Meyer and Wells, 1997).

In this report, the terms “soil(s)” refer to pedogenic soils formed in surficial deposits and possibly in sediments of the Dry Union Formation. Soil-horizon designations are those of the Soil Survey Staff (1999), Birkeland (1999), and Schoeneberger and others (2012). Colors of surficial deposits and sediments of the Dry Union Formation, followed by numbers and letters (for example, 10YR 7/2), are for air-dried samples of the matrix and were determined with the aid of Munsell soil color charts (Munsell Color, 1973). All other colors for sediments and bedrock in this report are field designations.

The approximate height above stream level of the top of gravelly stream alluvium of middle Pleistocene to Holocene age was determined with the aid of a 1:14,000-scale topographic map with 10-foot contour intervals, which was generated from lidar data. Height above stream level referred to in this report is the difference in altitude (expressed in meters) between stream level and the down-slope limit of gravelly stream alluvium at or near its depositional top.

In this report, the term “sericitized” refers to plagioclase that has been altered to the mineral sericite. The term “migmatite” refers to biotite gneiss containing numerous layers and irregular pods of light-gray or white granitic rock (leucosomes), typically 0.1–10 cm thick but locally thicker, which in some places are strongly folded along with the enclosing biotite gneiss. Formation of leucosomes is generally attributed to partial melting of the biotite gneiss during metamorphism (Johannes and Gupta, 1982; Olsen, 1982; Cole and Braddock, 2009).

**Table 1.** Particle-size analyses of loess and till in and near the Granite 7.5' quadrangle. Field-sample numbers and values from Shroba (1977).

[Sample locations are not shown on the geologic map and are not in the geologic database. Samples were treated with sodium peroxide ( $\text{Na}_2\text{O}_2$ ) to remove organic matter and with dilute hydrochloric acid (HCl) to remove any carbonate. Sodium hexametaphosphate ( $\text{NaPO}_3$ )<sub>6</sub> was then added as a dispersant to all samples, which were subjected to ultrasonic vibration for 15 minutes. Particle-size analyses were performed by sieve-and-pipette methods (Shroba, 1977). Values are fraction of total weight (percent).]

Deposit and field-sample number	Location (UTM)*	Percent < 2mm			Estimated percent >2 mm
		Sand <sup>1</sup> (2–0.05 mm)	Silt (0.05–0.002 mm)	Clay (<0.002 mm)	
<b>Loess</b>					
TL–6–2	0381100E 4321300N	21	52	27	10
<b>Youngest till<sup>2</sup></b>					
TL–5–3	0380140E 4320360N	72	21	7	80
<b>Till of Pinedale age younger unit</b>					
TL–6–5	0381100E 4321300N	60	32	8	50
TL–7–4	0381180E 4323950N	81	12	7	60
TL–8–1	0381200E 4324460N	58	30	12	40
TL–9–6	0382800E 4326800N	69	27	4	45
TL–10–6	0387100E 4325900N	72	24	4	60
TL–11–6	0388400E 4325600N	72	22	6	50
<b>Till of Pinedale age older unit</b>					
TL–12–6	0388040E 4324800N	77	18	5	70
<b>Till of pre-Bull Lake age</b>					
TL–14–7	0384490E 4329630N	76	13	11	70

<sup>1</sup>Size ranges are those of the Soil Survey Staff (1951).

<sup>2</sup>Youngest till in Boswell Gulch may be equivalent to deposits of the Satanta Peak advance of Benedict (1973b, 1985) in the Indian Peaks area of the Front Range.

\*Datum: North America 1927

## DESCRIPTION OF MAP UNITS

## Surficial Deposits

## Manmade Deposits

- af** **Artificial-fill deposits (latest Holocene)**—Engineered, compacted fill material composed of rock fragments and finer material was emplaced in the Twin Lakes and Mount Elbert Forebay earth-fill dams, above the penstocks connecting Mount Elbert Forebay and the Mount Elbert power plant, beneath parts of U.S. Highway 24 and Colorado Highway 82, and railroad tracks on the east side of the Arkansas River. Unit also locally includes small un-compacted deposits composed of rock fragments and finer material. Estimated thickness is 3–30 m
- pt** **Placer-tailings deposits (latest Holocene)**—Sandy, locally slightly bouldery, cobble gravel deposited by flowing water and power equipment. Unit also locally includes linear and irregularly shaped piles, as much as 3 m high, composed chiefly of cobbles and small boulders that appear to have been stacked by hand. These tailings piles locally contain boulders as large as 75 × 120 × 180 cm. Deposits of unit pt were first produced (about 1860–1883) by hand methods and ground sluicing and later (about 1884–1950) by hydraulic mining, dredging, and dragline in order to concentrate native gold (Parker, 1974a, 1974b). Placer-tailings deposits and gold were derived chiefly from cobbly glacial outwash (Qopy and Qoey) and fluvial gravel in the Dry Union Formation (Nd). Small tailings deposits between Cache Creek and Lake Creek were derived from deposits of alluvium and colluvium, undivided (Qac), colluvium (Qc), and fan deposits (Qf). A considerable amount of sand and finer sediment, transported by water flowing from areas of hydraulic mining along and near Cache Creek, was conveyed to the Arkansas River by a network of wooden flumes. From south to north, the most extensive placer-tailings deposits in the map area are along and near Cache Creek, Lake Creek, and Corske Creek. Estimated thickness is 1–10 m
- mt** **Mine-tailings deposits (latest Holocene)**—Four piles of un-compacted, very angular rock debris dumped near small abandoned mines. Two in Sheephead Gulch and Boswell Gulch near the southwest corner of the map area, and two near Twobit Gulch and Spring Creek east of the Arkansas River. Rock fragments in deposits of unit mt commonly range in size from about 50 cm in diameter to sand-sized particles. These deposits were produced during the exploration and (or) the extraction of gold-bearing ore, and locally silver-bearing ore, between the early 1860s and as late as the 1930s (Hedlund and others, 1983; Gese and Scott, 1993; Cappa and Bartos, 2007). Estimated thickness is 1–10 m

## Lacustrine Deposits

- Qb** **Beach deposits (latest Holocene)**—Mostly deposits of coarse sand, pebbly coarse sand, and pebble gravel. Deposits of unit Qb were formed, after the construction of the Twin Lakes dam in 1978, by wave action on outwash deposits of Pinedale age younger unit (Qopy) at two areas along the northern and eastern shores of the Twin Lakes Reservoir near an altitude of 2,805 m. Unit locally may include a minor amount of outwash of Pinedale age younger unit. Estimated thickness is 1–2 m; possibly thicker near the Twin Lakes dam (Tweto, 1978)

## Alluvial Deposits

- Qva** **Valley-floor alluvium (Holocene and latest Pleistocene?)**—Chiefly poorly sorted pebbly sand and pebbly to cobbly gravel in stream channels and floodplains and beneath low terraces (1–2 m high) along the Arkansas River and its major eastward flowing tributaries (table 2). Gravelly alluvium in floodplains is commonly overlain by about 30 cm or more of overbank sediment composed of slightly pebbly, very fine to very coarse sand that contains a minor amount of granules. Unit Qva locally includes bouldery deposits along tributary streams near the northern boundary of the map area, and small deposits of sheetwash alluvium (Qsw), small fan deposits (Qf), and colluvium (Qc) along valley margins. Unit Qva locally may

- include deposits of outwash of Pinedale age younger unit (**Qopy**). In glaciated valleys and in the valley of the Arkansas River, unit **Qva** probably overlies outwash of Pinedale age younger unit (**Qopy**). The long, low north-south trending fluvial scarp near the northern boundary of the map area corresponds with the western boundary of young anastomosing stream-channel and thin overbank sediments deposited by the Arkansas River. Deposits of unit **Qva** are a potential source of clean coarse aggregate. However, hydrologic factors such as locally high-water tables and periodic flooding, as well as geologic factors such as thickness of fine-grained overbank sediments and the amount of large cobbles and boulders, may render these deposits less suitable than some of the deposits of outwash gravel within the map area. Low-lying deposits are prone to periodic flooding. Estimated thickness is 5–10 m
- Qsw** **Sheetwash alluvium (Holocene and late Pleistocene)**—Few small deposits commonly composed of slightly pebbly to pebbly, slightly silty, fine to very coarse sand and a minor amount of granules deposited by overland flow on gentle slopes along and near the Arkansas River. Much of the unit is probably derived from erosion of sandy layers in the Dry Union Formation (**Nd**) and *grus* formed chiefly by chemical weathering and disintegration of the granite of Granite (**Ygrg**). Charcoal fragments collected from the base of a thin (29 cm) deposit of sheetwash alluvium that overlies outwash gravel of pre-Bull Lake age about 2 km north of the map area yielded a radiocarbon age of about 4.0 ka (Nelson and Shroba, 1998). Thin deposits of sheetwash alluvium, too narrow to map separately, locally overlie outwash deposits at and near valley sides. Low-lying areas of unit **Qsw** are susceptible to sheet flooding due to unconfined overland flow, and locally may be susceptible to gullyng. Estimated thickness is 1–7 m
- Qsw/Nd** **Sheetwash alluvium over Dry Union Formation**
- Outwash of Pinedale age (late Pleistocene)**—Glaciofluvial deposits composed chiefly of clast-supported, poorly sorted, and crudely stratified coarse cobbly pebble gravel deposited by meltwater streams during the Pinedale glaciation. Units (**Qopy** and **Qopo**) locally include layers and lenses of sandy pebble gravel and pebbly sand. Gravelly deposits are locally overlain by about 30–50 cm of overbank sediment composed of pebbly sand (Nelson and Shroba, 1998). Deposits proximal to till of Pinedale age (**Qtpy** and **Qtpo**) are bouldery and locally contain boulders as large as 110 × 140 × 180 cm. Soils formed in outwash of Pinedale age have cambic B (**Bw**) or weakly developed argillic B (**Btj** and **Bt**) horizons about 30 cm thick (Nelson and Shroba, 1998). The exposed portions of most surface boulders of granitic composition show little evidence of weathering. Outwash deposits of Pinedale age are a potential source of clean coarse aggregate, although locally they contain a significant amount of large cobbles and boulders and some of the biotite-rich pebbles and small cobbles are partly or completely disintegrated, probably due to chemical weathering and expansion of biotite (Wahrhaftig, 1965; Isherwood and Street, 1976)
- Qopy** **Younger unit**—Unit underlies terraces and terrace remnants about 2–3 m above Clear Creek, Cache Creek, and Corske Creek, 2–7 m above Lake Creek, and 2–12 m above the Arkansas River (table 2). Unit along Lake Creek near Balltown locally contains boulders as large as about 270 × 300 × 300 cm. These boulders probably were deposited by one or more glacial-outburst floods (*jökulhlaups*) initiated by the rapid drainage of a lake or lakes impounded by glacial ice or moraines of unit **Qtpy** in the vicinity of the Twin Lakes dam. Unit **Qopy** includes one small kame deposit composed chiefly of sandy pebble gravel, about 0.7 km southwest of the Mt. Elbert power plant. About 0.5 km southeast of Balltown (fig. 2), unit **Qopy** locally includes placer-tailings deposits (**pt**) that are too small to show at map scale. In the valleys of Lake Creek and Clear Creek, deposits of unit **Qopy** are graded to and coeval with till of unit **Qtpy**, which dates from about 22–15 ka (D.P. Dethier, Williams College, Williamstown, Mass., written commun., 2010; Young and others, 2011). Estimated thickness is 3–20 m; possibly 50 m or more on the up-valley sides of terminal moraines in the valleys of Lake Creek and Clear Creek
- Qopo** **Older unit**—Unit underlies terrace remnants about 18 m above Lake Creek and 18–24 m above the Arkansas River (table 2). On the east side of the Arkansas River, the top part of the unit locally contains many very angular rockfall-derived boulders as large as 140 × 160 × 290

cm. The top part of the unit at the eastern boundary of the map area southeast of the town of Granite contains boulders as large as  $140 \times 255 \times 520$  cm, and may be a glacial-outburst flood deposit. Near the confluence of Lake Creek with the Arkansas River, deposits of unit **Qopo** are graded to and coeval with till of unit **Qtpo**, which is older than 20 ka in the valley of Lake Creek (D.P. Dethier, Williams College, Williamstown, Mass., written commun., 2010) and older than 22 ka in the valley of Clear Creek (Young and others, 2011). Estimated thickness is 3–20 m

**Outwash of Bull Lake age (late and middle Pleistocene)**—Glaciofluvial deposits composed chiefly of coarse cobbly pebble gravel deposited by meltwater streams during the Bull Lake glaciation. Units (**Qoby** and **Qobo**) locally include layers and lenses of sandy pebble gravel and pebbly sand. Gravelly deposits are locally overlain by about 20–30 cm of overbank sediment composed of pebbly sand (Nelson and Shroba, 1998). Deposits proximal to till of Bull Lake age (**Qtb**) are bouldery and locally contain boulders as large as  $120 \times 130 \times 260$  cm. Soils formed in outwash of Bull Lake age have argillic B (Bt) horizons commonly about 40–65 cm thick (Nelson and Shroba, 1998). The exposed portions of some surface boulders of granitic composition are extensively weathered. Some of these boulders are “hat shaped” and have “brims” about 15–60 cm wide. Outwash deposits of Bull Lake age are a potential source of coarse aggregate, although locally they contain a significant amount of large cobbles and boulders and many of the biotite-rich pebbles and cobbles are partly or completely disintegrated

**Qoby** **Younger unit (late and middle Pleistocene)**—Unit underlies terrace remnants about 37 m above the Arkansas River (table 2). On the south side of Lake Creek, deposits of unit **Qoby** are graded to and coeval with till of unit **Qtb**. On the north side of Lake Creek, at the confluence of Lake Creek and the Arkansas River, the top of unit **Qoby** locally is overlain by scattered boulders of granitic composition that are commonly 1–2 m long. The largest boulder is  $>100 \times 210 \times 400$  cm. These boulders probably were deposited by one or more glacial-outburst floods initiated by the rapid drainage of a lake or lakes impounded by glacial ice or moraines of unit **Qtb** farther upstream in the Lake Creek drainage. Estimated thickness is 3–20 m

**Qobo** **Older unit (middle Pleistocene)**—Unit underlies terrace remnants about 43 m above the Arkansas River (table 2). On the south side of Lake Creek, deposits of unit **Qobo** are graded to and coeval with till of unit **Qtb**. Estimated thickness is 3–20 m

**Outwash of pre-Bull Lake age (middle Pleistocene)**—Glaciofluvial deposits composed chiefly of coarse cobbly pebble gravel deposited by meltwater streams probably during two or more pre-Bull Lake glaciations, and one deposit of pebble gravel of non-glacial origin. Units (**Qoey** and **Qo eo**) locally include layers and lenses of sandy pebble gravel and pebbly sand. Gravelly deposits are locally overlain by about 20–30 cm of slightly pebbly, slightly silty sand. These thin layers may be overbank sediment or loess (windblown silt) mixed with underlying alluvium (Nelson and Shroba, 1998). Deposits proximal to till of pre-Bull Lake age (**Qte**) are bouldery and locally contain boulders as large as  $>80 \times 190 \times 235$  cm. Soils formed in outwash of pre-Bull Lake age have clayey argillic B (Bt) horizons that are 30 cm to greater than 120 cm thick and carbonate-enriched calcic horizons (Bk and Btk) horizons 45 cm to greater than 60 cm thick with stage I to strong stage II carbonate morphology (Alan R. Nelson, Bureau of Reclamation, Denver, Colo., written commun., 1981). The exposed portions of surface boulders of granitic composition are extensively weathered. Some of these boulders are “hat shaped” and have “brims” about 50–160 cm wide. Outwash deposits of pre-Bull Lake age are a potential source of coarse aggregate, although locally they contain a significant amount of large cobbles and boulders and all of the biotite-rich pebbles and cobbles are partly or completely disintegrated

**Qoey** **Younger unit**—Unit underlies terrace remnants at three levels about 55–61, 67–72, and 79–87 m above the Arkansas River (table 2). Unit also includes a deposit of pebble gravel rich in granitic clasts on the north side of Sawmill Gulch, about 80 m above the Arkansas River, near the northern boundary of the map area. About 1 km northwest of the Granite Cemetery, the uppermost deposits of unit **Qoey** includes a S. 50° E.-trending flood-

deposited gravel body about 40 m long and 20 m wide that contains widely spaced boulders. Most of the boulders are granitic in composition and are commonly 40–120 cm long. Boulders on unit **Qoey** transported by the Arkansas River are commonly about 30–40 cm in length and are locally as large as  $70 \times 75 \times 120$  cm. West of the river boulders transported by tributary streams are slightly larger. Commonly, they are about 30–60 cm in length and are locally as large as  $\geq 85 \times 120 \times 200$  cm. On the north side of Clear Creek, the highest deposits of unit **Qoey** appear to be graded to and may be coeval with till of unit **Qte**. Estimated thickness is 1.5–15 m for deposits east of the Arkansas River and 10–20 m for deposits west of the Arkansas River

**Qofo** **Older unit**—Unit underlies a high terrace north of Lake Creek and includes two small gravel deposits east of the Arkansas River. The tops of the terrace and the gravel deposits are about 95–110 m above the Arkansas River (table 2). Clast composition indicates that eastward flowing streams tributary to the Arkansas River as well as the southeast-flowing Arkansas River deposited the glaciofluvial gravel beneath the large northwest-southeast-trending terrace; whereas, the two small gravel deposits east of the Arkansas River were deposited by the Arkansas River. The uppermost deposits of unit **Qofo** at the south end of the high terrace, just north of Lake Creek, include two parallel S. 50° E. trending flood-deposited gravel bodies about 80 m long and 20 m wide that contain many large boulders. These boulders are granitic in composition and are commonly 140–240 cm long. The largest boulder is  $\geq 180 \times 200 \times 420$  cm. These boulders probably were deposited by one or more glacial-outburst floods, which may have originated in the Lake Creek drainage. Deposits of unit **Qofo** on the north side of Lake Creek are graded to and coeval with till of unit **Qte**. Unit **Qofo** may be correlative with outwash gravel 4 km southwest of Nathrop, Colo. (Nelson and Shroba, 1998, p. 354, fig. 3). The gravel near Nathrop is about 100 m above the Arkansas River and is overlain by fine-grained sediment that contains Lava Creek B ash (Scott and others, 1975; Izett and Wilcox, 1982), which is about 640 ka (Lanphere and others, 2002). Unit **Qofo** is likely to be younger than the next older outwash gravel located 7 km southwest of Nathrop. This older gravel is about 165 m above the Arkansas River and is overlain by fine-grained sediment that contains the Bishop ash (Scott and others, 1975; Izett and others, 1988), which is about 760 ka (Sarna-Wojcicki and others, 2000). Estimated thickness is 2–5 m for the deposits east of the Arkansas River and 10–20 m for the deposit west of the Arkansas River

#### Alluvial and Mass-movement Deposits

**Qac** **Alluvium and colluvium, undivided (Holocene to middle? Pleistocene)**—Composed chiefly of undifferentiated sandy and gravelly stream alluvium, sheetwash alluvium (**Qsw**), fan deposits (**Qf**), colluvium (**Qc**), and locally other mass-movement deposits along minor streams and on adjacent lower (toe) slopes. Along Cache Creek, the top 50–100 cm of unit **Qac** consists of brownish-black, humic overbank sediments composed of slightly pebbly and locally granule-rich silty sand. Sediments at the mouth of Bartlett Gulch, near the western boundary of the map area, probably include deposits of humic silty sand. Unit **Qac** along Cache Creek and near the head of Lost Canyon, about 5 km south of the Twin Lakes Reservoir, locally contains small placer-tailings deposits (**pt**). Low-lying areas of unit **Qac** adjacent to stream channels may be subject to periodic stream flooding and debris-flow deposition. The adjacent toe slopes may be susceptible to local sheet flooding due to unconfined overland flow, and locally to gullying. Estimated thickness is 3–15 m

**Qf** **Fan deposits (Holocene to middle? Pleistocene)**—Mostly poorly sorted, clast-supported, slightly bouldery pebble and cobble gravel and locally minor deposits composed of matrix-supported, pebbly and cobbly silty sand. Clasts are commonly very angular to subangular. Boulders are locally 1–2 m long. Deposited chiefly by streams and debris flows in fan- and tongue-shaped accumulations near base of moderate to steep slopes of the Sawatch Range and locally along the steep flanks of lateral moraines of unit **Qtpy** in the valleys of Lake Creek and Clear Creek. Snails (cf. *Catinella*) from a fan deposit along the north shore of Clear Creek Reservoir (NW¼ of sec. 7, T. 12 N., R. 79 W.) had amino

acid ratios that suggest a pre-Holocene age for some beds within the fan (A.R. Nelson, U.S. Geological Survey (USGS), written commun., 2013). Unit probably locally includes deposits of sheetwash alluvium (Qsw), colluvium (Qc), and probably hyperconcentrated-flow deposits. Low-lying areas on the unit may be subject to periodic stream flooding and debris-flow deposition. Estimated thickness is 1–15 m

#### Mass-movement Deposits

- Qta** **Talus deposits (Holocene to middle? Pleistocene)**—Very angular and angular boulders as much as 5 m long and smaller rock fragments deposited chiefly by rock and snow avalanche, rockfall, rock slide, and debris flow at the base of cliffs and steep slopes where debris forms broad aprons, small cones, and fan-shaped deposits. Unit locally includes rubbly scree deposits and small debris-flow deposits. Unit **Qta** near the head of Boswell Gulch, near southwest corner of the map area, locally overlies rock-glacier deposits (**Qrg**) and till of Pinedale age younger unit (**Qtpy**). A thin (10 cm) layer, that may be loess mixed with or infiltrated into the underlying debris, locally overlies unit **Qta** near the head of Boswell Gulch. Talus deposits in the valley of Clear Creek, which extend below the up-slope limit of the glacial ice extent during the Pinedale glaciation, postdate the retreat of Pinedale ice. Deposits derived from steep Proterozoic bedrock outcrops along the Arkansas River, too small to show at map scale, locally overlie deposits of unit **Qopo**. Estimated thickness is 1–15 m
- Qc** **Colluvium (Holocene to middle? Pleistocene)**—Non-sorted, mostly matrix-supported deposits that consist of clay-, silt-, and sand-sized particles, and clasts that range in size from granules to large boulders. Deposits of unit **Qc** are formed by a variety of mass-movement processes. Unit includes deposits on sagebrush-covered and forested slopes as well as deposits in the large alpine area, at and above present day treeline (about 3,480–3,690 m) near the head of Lost Canyon, where freeze-thaw action has promoted mass-movement and downslope transport. The composition of the deposits reflects that of the underlying or nearby bedrock or sediments from which the colluvium was derived. In areas of Proterozoic bedrock, unit **Qc** probably consists chiefly of transported grus and corestones. Unit **Qc** locally includes material transported by frost creep, solifluction, nivation, periglacial processes, sheetwash, sliding, debris flow, hyperconcentrated flow, earth flow, and rockfall, and by minor ephemeral streams. In the large alpine area near the head of Lost Canyon, unit **Qc** locally includes solifluction deposits that form large stair-steps with broad treads and low risers that are roughly parallel to the contours. These deposits are probably similar in origin to the turf-bank terrace deposits (Benedict, 1970) in the Indian Peaks region of the Front Range, about 120 km northeast of the map area. Unit **Qc** locally includes small bodies of Proterozoic bedrock too small to show at map scale. Estimated thickness is 3–15 m
- Qc/Nd** **Colluvium over Dry Union Formation**
- Qls** **Landslide deposits (Holocene to middle? Pleistocene)**—Deposits composed of unsorted and unstratified debris of various sizes commonly at or near the base of moderate to steep slopes. These deposits commonly have hummocky topography. The largest landslides and landslide deposits form on unstable slopes that are underlain by sediments of Dry Union Formation (**Nd**) near the northern and western boundary of the map area and on highly fractured Proterozoic bedrock (**Ygrg**) near the southwest corner of the map area. Younger landslide deposits are commonly bounded upslope by crescent-shaped headwall scarps and downslope by lobate toes. Older landslide deposits commonly lack scarps and have more muted surface morphology. Unit **Qls** includes material displaced chiefly by rotational debris slides, translational rock and debris slides, debris flows, and possibly minor earth flows as defined by Varnes and Cruden (1996). The sizes and lithologies of the clasts and the particle-size distributions of the matrices of these deposits reflect those of the displaced bedrock units and surficial deposits. The large landslide deposits on the north side of Cache Creek, on the north side of the Twin Lakes Reservoir, at the northwest corner of the map area, and on the west side of the Arkansas River near the northern boundary of the

map area are derived chiefly from sediments of unit Nd. Locally, these landslide deposits contained small intact blocks and deformed lens-shaped bodies of unit Nd. Unit Qls locally includes small deposits of sheetwash alluvium (Qsw) and marsh or bog deposits that accumulated in unfilled depressions. The large landslide deposit derived from Proterozoic bedrock north of Clear Creek near the southwest corner of the map area locally has prominent up-slope facing scarps, referred to as sackungen (Varnes and others, 1989; McCalpin and Irvine, 1995). Small talus deposits (Qta), too small to show at map scale, locally mantle the large bedrock-derived landslide deposit at the southern boundary of the map area. The landslide deposits along the north and south shores of Twin Lakes Reservoir, along the valley of Clear Creek, and near the northwest corner of the map area formed after melting and retreat of the Pinedale glaciers in these areas. Some of these latter landslide failures of unconsolidated or weakly consolidated sediments on steep slopes may have been due in part to steep slopes, high water content, and high pore-water pressures in unstable sediments as well as the removal of lateral support formerly provided by glacial ice prior to melting. Landslide deposits may be prone to continued movement or reactivation due to both natural and human-induced processes. Estimated thickness is 5–45 m

#### Glacial Deposits

- Qrg** **Rock-glacier deposits (early Holocene? and latest Pleistocene?)**—Large masses composed of boulder- to silt-sized rock debris that have relatively steep fronts and flanks, and gently sloping upper surfaces. Deposits of unit Qrg mantle lobate (width greater than length) rock glaciers, as defined by Wahrhaftig and Cox (1959). They formed at or near the base of talus deposits (Qta) at an elevation of 3,500–3,710 m in the cirque at the head of Boswell Gulch, near the southwest corner of the map area. Deposits of unit Qrg probably consist of a veneer of angular boulder- and cobble-sized clasts that overlies smaller debris. Rock fragments on and within unit Qrg probably accumulated chiefly by rockfall and locally by sliding, and snow and rock avalanche. These fragments are probably moved downslope by creep or flowage. The upper surfaces of the unit locally have prominent longitudinal and transverse ridges and furrows, and a few collapse pits. The lower part of the unit may be ice-cemented debris or may have an ice core. Lobate rock glaciers commonly are ice-cemented (Benedict, 1973a; White, 1976), and likely formed under periglacial conditions (Barsch, 1987). Unit Qrg overlies till of Pinedale age younger unit (Qtpy), and locally may include minor debris-flow deposits and talus deposits (Qta) displaced by creep. A thin (5–10 cm) loess mantle locally overlies rock-glacier deposits in La Plata Basin (Shroba, 1977), about 8 km west of the map area. Many of the rock-glacier deposits in Colorado are considered to be of latest Pleistocene or early Holocene age (Meierding and Birkeland, 1980). Estimated maximum thickness is 45 m
- Till of Pinedale age (late Pleistocene)**—Mostly non-sorted and non-stratified, subangular and subrounded boulders to granules in a matrix composed of very slightly clayey, slightly silty sand (table 1). Deposits locally may contain significant amounts of silt, sand, or gravel in areas where glaciers have eroded pro-glacial lacustrine sediments, outwash, or sediments of the Dry Union Formation (Nd). Material >2 mm in diameter is estimated to be 40–70 percent of the unit. Some of the surface boulders of granitic composition are as large as  $\geq 450 \times 550 \times 700$  cm. Units (Qtpy and Qtpo) locally include small bodies of silt, sand, and gravel deposited by meltwater, and are locally overlain by till-derived landslide deposits (Qls). Soils formed in till of Pinedale age have thin (7–22 cm) cambic B (Bw) horizons, or they lack B horizons and have A/Cox or A/Bk profiles (Shroba, 1977; Nelson and Shroba, 1998). Calcic (Bk) horizons have weak stage I carbonate morphology. Depth of oxidation (base of Cox horizon) ranges from about 50 to 170 cm. Unoxidized till is gray (commonly 10YR 7/2, 2.5Y 7/2, and 5Y 7/2), whereas soil B horizons are light brown (7.5 YR 5/3). Locally, as much as 15 percent of the biotite-rich pebbles and small cobbles within soils are partly or completely disintegrated (Shroba, 1977). The exposed portions of some of the surface boulders of granitic composition are weathered. Probably less than 1 percent of these weathered boulders are “hat shaped” and have “brims” about 10–25 cm wide. Surface boulders typically are more abundant on moraines of the Pinedale glaciation than on those

of older glaciations. Deposits form bouldery, prominent, sharp-crested lateral moraines along Clear Creek as well as lateral and end moraines with similar surface morphology along Lake Creek and in Boswell Gulch. Moraines locally have unfilled and undrained depressions. Pinedale glaciers were as thick as 350 m in the valley of Lake Creek and flowed 31 km down valley from an altitude of about 3,900 m to an altitude of about 2,780 m. In the valley of Clear Creek, glaciers were as thick as 400 m and flowed 23 km down valley from an altitude of about 3,800 m to an altitude of about 2,680 m. Estimated thickness is 1.5–45 m

- Qtpy** **Younger unit**—In the valleys of Lake Creek and Clear Creek, till of unit **Qtpy** is graded to and coeval with outwash deposits of unit **Qopy**. Six  $^{10}\text{Be}$  surface-exposure ages of boulders on moraines within the map area indicate that till of unit **Qtpy** in the valleys of Lake Creek and Clear Creek dates from about 22–15 ka (D.P. Dethier, Williams College, Williamstown, Mass., written commun., 2010; Young and others, 2011). These ages are compatible with five  $^{10}\text{Be}$  surface-exposure ages (21–16 ka) reported by Brugger (2007) for boulders on deposits that form the outermost moraines of the last glacial maximum in the Taylor River drainage, about 20 km southwest of the map area. There are six  $^{10}\text{Be}$  surface-exposure ages of glacially scoured bedrock surfaces. Three are located in the valley of Lake Creek about 3, 8, and 17 km west of the map area, one in the valley of Clear Creek near the southern boundary of the map area, and two  $^{10}\text{Be}$  surface-exposure ages within a cirque about 12 km northwest of the map area. These ages suggest that much or all of unit **Qtpy** in Boswell Gulch dates from about 15–13 ka (Mason and others, 2011; Young and others, 2011; Ruleman and others, 2013). Unit **Qtpy** in Boswell Gulch is down valley of and older than bouldery till that forms a small sharp-crested moraine at an altitude of 3,780 m at the head of Boswell Gulch, about 0.7 km west of the map area. This till is the youngest till observed in the cirques in and near the map area (sample TL–5–3, table 1). It may be correlative with deposits of the Satanta Peak advance of Benedict (1973b, 1985) in the Indian Peaks region of the Front Range, about 120 km northeast of the map area. Deposits of the Satanta Peak advance have minimum-limiting radiocarbon ( $^{14}\text{C}$ ) ages of about 9.9 and 9.7 ka (Benedict, 1973b). These deposits probably date from 12 to 10 ka B.P. (Benedict, 1985), and may be coeval with glacial deposits formed during the Younger Dryas climatic event (about 11–10 ka) in northwestern North America and northern Europe (Menounos and Reasoner, 1997). A thin loess mantle, about 20 cm thick, locally overlies unit **Qtpy** near the head of Boswell Gulch (sample TL–6–2, table 1). Loess was not observed elsewhere on unit **Qtpy** or on any of the other till units within the map area
- Qtpo** **Older unit**—Near the confluence of Lake Creek with the Arkansas River, the till of unit **Qtpo** is graded to and coeval with outwash deposits of unit **Qopo**. The oldest of five  $^{10}\text{Be}$  surface-exposure ages for boulders on the outer lateral moraines of unit **Qtpy** on the north side of Clear Creek Reservoir suggest that nearby deposits of unit **Qtpo** are older than 22 ka (Young and others, 2011). Unit **Qtpo** may be about 30 ka (Nelson and others, 1979). It may be correlative with an early advance of Pinedale ice about 47–34 ka in the northern Yellowstone area (Sturchio and others, 1994) or with an earliest Pinedale advance at about 70–50 ka at McCall, Idaho, and in the Wind River Range near Lander, Wyo. (Dahms, 2004; Pierce and others, 2011)
- Qtb** **Till of Bull Lake age (late and middle Pleistocene)**—Mostly non-sorted and non-stratified, subangular and subrounded boulders to granules in a matrix composed of very slightly clayey, slightly silty sand. Deposits locally may contain significant amounts of silt, sand, or gravel in areas where glaciers have eroded proglacial lacustrine sediments, outwash, or sediments of the Dry Union Formation (Nd). Material >2 mm in diameter is estimated to be 40–70 percent of the unit. Some of the surface boulders of granitic composition are as large as  $\geq 140 \times 350 \times 380$  cm. Unit locally includes small deposits of sheetwash alluvium (**Qsw**) and colluvium (**Qc**) on the lower flanks of moraines and probably locally includes small bodies of silt, sand, and gravel deposited by meltwater. Soils formed in till of Bull Lake age have argillic B (Bt) horizons about 45–50 cm thick. These horizons are about two to three times as thick and contain about twice as much clay compared to argillic B horizons formed in till of Pinedale age. Many of the biotite-rich pebbles and cobbles

within soils formed in till of Bull Lake age are partly or completely disintegrated. The exposed portions of many of the surface boulders of granitic composition are extensively weathered. Some of these boulders are “hat shaped” and have “brims” about 20–130 cm wide. Surface boulders typically are less abundant and are more weathered on moraines of the Bull Lake glaciation than on those of the Pinedale glaciation. Deposits of unit **Qtb** form prominent lateral moraines that have rounded crests along and near the Twin Lakes Reservoir, Clear Creek Reservoir, and Dry Columbia Gulch, near the southeast corner of the map area. Undrained depressions are rare. K-Ar and  $^{230}\text{Th}/\text{U}$  analyses that constrain the ages of glaciofluvial deposits near the type area for deposits of the Bull Lake glaciation along the northern flank of the Wind River Range, Wyo., and ages of glacial deposits near West Yellowstone indicate that the Bull Lake glaciation probably began prior to  $167 \pm 6.4$  ka (possibly 190 ka) and may have continued until about  $122 \pm 10$  ka (Sharp and others, 2003; Pierce, 2004).  $^{10}\text{Be}$  ages for the West Yellowstone glacial system yield a mean age of  $136 \pm 13$  ka and oldest ages of about 151–157 ka (Licciardi and Pierce, 2008).  $^{10}\text{Be}$  and  $^{26}\text{Al}$  analyses of surface boulders on moraines composed of till of Bull Lake age near Nederland, Colo., about 110 km northeast of the map area, yielded minimum age estimates of  $101 \pm 21$  and  $122 \pm 26$  ka (Schildgen and others, 2002). These age estimates are in accord with a uranium-trend age estimate of  $130 \pm 40$  ka for till of Bull Lake age (Shroba and others, 1983) near Allens Park, Colo., about 135 km northeast of the map area. On the south side of Lake Creek, till of unit **Qtb** is graded to and coeval with outwash deposits of units **Qoby** and **Qobo**. Glacial ice during the Bull Lake glaciation flowed 31 km down the Lake Creek valley to an altitude of about 2,800 m, and 21 km or more down the valley of Clear Creek to an altitude of about 2,780 m. Estimated thickness is 1.5–30 m

Qte

**Till of pre-Bull Lake age (middle Pleistocene)**—Mostly non-sorted and non-stratified, subangular and subrounded boulders to granules in a matrix composed of slightly clayey, slightly silty sand (sample TL-14-7, table 1). Deposits locally may contain significant amounts of silt, sand, or gravel in areas where glaciers have eroded pro-glacial lacustrine sediments, outwash, or sediments of the Dry Union Formation (Nd). Material  $>2$  mm in diameter is estimated to be 40–70 percent of the unit. Some of the surface boulders of granitic composition are as large as  $\geq 125 \times \geq 190 \times 210$  cm. North of Lake Creek, unit locally includes till assigned to the Bull Lake glaciation by Nelson and Shroba (1998). Unit probably locally includes small deposits of sheetwash alluvium (**Qsw**) and colluvium (**Qc**) on the lower flanks of moraines as well as small bodies of silt, sand, and gravel deposited by meltwater. Soils formed in till of pre-Bull Lake age have clayey argillic B (Bt) horizons about 90 cm thick that contain about twice as much clay as those formed in till of Bull Lake age (Shroba, 1977; Nelson and Shroba, 1998). All of the biotite-rich pebbles and cobbles within soils are partly or completely disintegrated. The exposed portions of many of the surface boulders of granitic composition are extensively weathered. Some of these boulders are “hat shaped” and have “brims” about 95–105 cm wide. Surface boulders typically are much less abundant on till of pre-Bull Lake age than on those of younger glaciations. This difference in relative abundance is probably due in part to weathering of surface boulders and possibly to fewer boulders in till of pre-Bull Lake age. Deposits of unit **Qte** form subdued lateral and end moraines that have broad rounded crests on the north side of the Twin Lakes Reservoir; those on the north side of Clear Creek lack morainal morphology. Undrained depressions were not observed. Till of unit **Qte** on the north side of Lake Creek is graded to and coeval with outwash deposits of unit **Qoeo**, the top of which is 95–110 m above the Arkansas River. Till of unit **Qte** on the north side of Clear Creek appears to be graded to and may be coeval with the highest outwash deposits of unit **Qoey**, the top of which is 79–87 m above the Arkansas River. If deposits of unit **Qoeo** on the north side of Lake Creek are correlative with gravelly alluvium 98 m above the Arkansas River near Nathrop, Colo., which is overlain by Lava Creek B ash, then deposits of unit **Qte** on the north side of Lake Creek are correlative with till of the Sacagawea Ridge glaciation. Deposits of this glaciation are located on the northern flank of the Wind River Range, Wyo., where the uppermost part of outwash of the Sacagawea Ridge glaciation contains water-lain Lava Creek B tephra (Richmond, 1976; Izett and Wilcox, 1982; Richmond and Murphy, 1989). The stratigraphic position of this tephra layer suggests that the Sacagawea Ridge glaciation occurred during marine

oxygen isotope stage 16 (Chadwick and others, 1997), about 675–621 ka (Lisiecki and Raymo, 2005). The height of the highest outwash gravel of unit **Qoey**, relative to that of unit **Qoee**, suggests that deposits of unit **Qte** on the north side of Clear Creek may have accumulated during a younger pre-Bull Lake glaciation that occurred after about 620 ka. Glaciers during one or more pre-Bull Lake glaciations flowed 31 km down the Lake Creek valley to an altitude of about 2,800 m and 21 km or more down the valley of Clear Creek to an altitude of about 2,780 m. Estimated thickness is 1.5–30 m

### **Bedrock Units**

#### **Post-volcanic Sediments**

**Nd** **Dry Union Formation (lower Pliocene? and Miocene)**—Deposits of the Dry Union Formation within the map area are exposed in small widely spaced outcrops and manmade cuts about 0.5–10 m high that range in altitude from about 2,780 to 2,940 m. Unit consists chiefly of fluvial and fan deposits. Fluvial deposits probably were transported and deposited by perennial, through-flowing streams. Fan deposits consist chiefly of stream-channel alluvium of minor tributary streams (referred to hereafter as stream-channel alluvium), sheetflood alluvium, and debris-flow deposits. These fan deposits were deposited by ephemeral and intermittent streams, sheetfloods, and debris flows, respectively. Fluvial deposits are locally exposed in an area about 2 km wide and 7 km long that is parallel to and just west of the canyon cut in Proterozoic basement rocks by the Arkansas River. Fan deposits composed of stream-channel alluvium, sheetflood alluvium, and debris-flow deposits are locally exposed east of the Arkansas River north of Twobit Gulch as well as west of the Arkansas River at several localities. Deposits of unit **Nd** commonly are unconsolidated to moderately consolidated, non-sorted or poorly sorted, and crudely stratified to well stratified. Gravel is commonly clast supported and is locally matrix supported. Biotite-bearing clasts are partly or completely disintegrated. Fluvial deposits of unit **Nd** north of Balltown consist chiefly of massive beds of coarse cobble gravel about 30 m thick that overlies finer-grained, lower-energy, fluvial deposits at least 60 m thick. The underlying exposed lower-energy fluvial deposits of unit **Nd** consist chiefly of pebbly sand, sandy pebble gravel, and slightly cobbly pebble gravel in lenses and layers about 3–200 cm thick. Fluvial deposits near the Granite Cemetery are at least 24 m thick. They consist chiefly of cobble gravel, cobbly pebble gravel, sandy pebble gravel, a subordinate amount of pebbly sand, and a minor amount of sandy silt in lenses and layers about 20–300 cm thick. Fluvial deposits, exposed in a small outcrop about 550 m south-southwest of the Granite Cemetery, overlie debris-flow(?) deposits of unit **Nd**. Orientations of imbricated cobbles observed in one exposure, just west of the Granite Cemetery, suggest fluvial transport towards the southeast. Clasts in fluvial deposits are angular to subrounded. Stream-channel alluvium commonly consists of deposits of sand and gravel in lenses and layers of sand about 5–250 cm thick. Sand bodies locally contain lenses of pebble gravel, and gravel bodies locally contain lenses of pebbly sand. Stream-channel alluvium locally includes small channel-fill deposits about 30–210 cm thick. Two small channel-fill deposits, observed near the base of unit **Nd** north of Twobit Gulch, consist of small-pebble conglomerate that is weakly cemented by secondary calcium carbonate. Clasts commonly are angular. Sheetflood alluvium commonly consists of alternating layers of mud and layers of pebble gravel or pebbly sand. Muddy upper layers are about 1–15 cm thick and pebbly lower layers are about 5–25 cm thick. These deposits form depositional packages about 120–250 cm in total thickness. Debris-flow deposits commonly consist of layers of pebbly sand above layers of cobbly, sandy pebble gravel. Sandy upper layers are about 15–50 cm thick and gravelly lower layers are about 50–200 cm thick. Clasts in debris-flow deposits commonly are very angular and subangular. Original thickness of unit **Nd** is unknown, because the unit was deposited on an irregular surface and the top of the unit is eroded (Tweto, 1961). Thickness based on outcrop and drill-hole data (Tweto, 1978; D.A. Ostenaar, S.L. Losh, and A.R. Nelson, Bureau of Reclamation, Denver, Colo., written commun., 1981) is at least 240 m on the north side of the Twin Lakes Reservoir. Preserved thickness in the subsurface, about 11 km north of the map area, may be as much as 1,200 m (Tweto and Case, 1972)

## Volcanic Rocks

- F<sub>e</sub>va** **Vesicular andesite flow (upper Oligocene)**—Pinkish-brown, finely vesicular, massive andesite porphyry south of Twin Lakes Reservoir. Consists of about 25 percent plagioclase, both as small microlites and large (to 5 mm long) crystals, 6 percent augite, 2 percent hornblende, and 2 percent opaque minerals. Glassy matrix composes about 65 percent of rock. Previously mapped as Grizzly Peak Tuff by Fridrich and others (1998); however, no rhyolite tuff was observed, only rubbly blocks are exposed in two small areas about 1.7 km southwest of the Twin Lakes dam. New <sup>40</sup>Ar/<sup>39</sup>Ar age on hornblende is 24.12±0.12 Ma (sample GR11–711, table 3)

## Hypabyssal Rocks

At least three suites of felsic intrusive rocks are recognized within the map area: (1) ~65 Ma rhyolite dikes, (2) ~40 Ma felsic plutons and dikes, and (3) a ~24 Ma rhyolite stock and inferred coeval microtonalite plug. Each is described below.

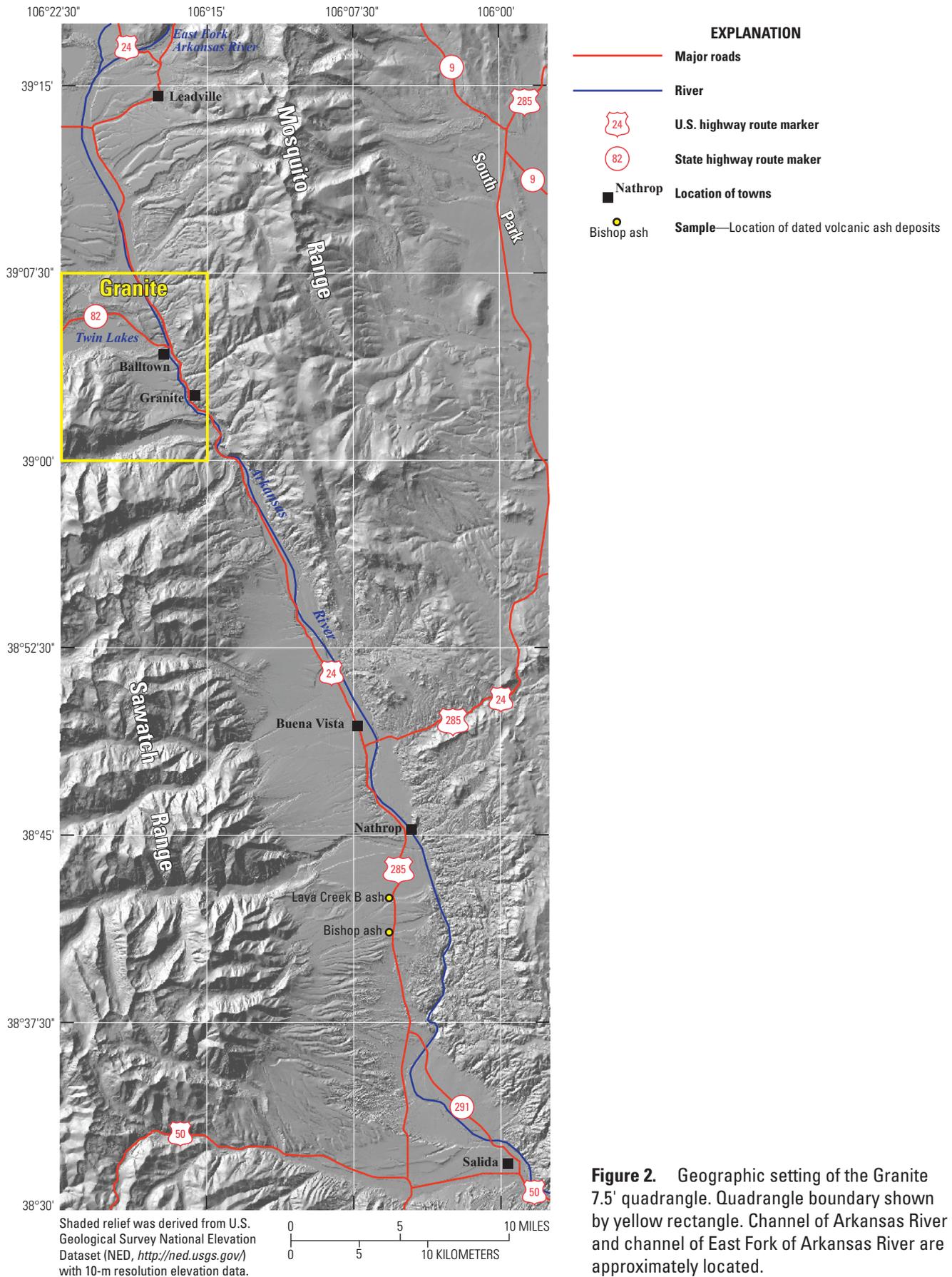
- F<sub>e</sub>rc** **Hypabyssal rhyolite of Clear Creek (late Oligocene)**—Pale pinkish-gray, porphyritic rhyolite(?) poorly exposed on ridge near southeast corner of the map area. Composed of about 55 percent glass with abundant quartz and feldspar microlites. Phenocrysts as long as 3 mm are composed of about 25 percent fragmental quartz, 9 percent plagioclase, 10 percent sanidine, 1 percent biotite, and trace opaque minerals. Contains many rounded fragments and clasts of Proterozoic gneiss as long as 1.5 cm. Locally contains small vugs. Rhyolitic composition inferred from phenocryst ratios. New <sup>40</sup>Ar/<sup>39</sup>Ar age on feldspar (sanidine?) is 24.3±0.6 Ma (sample GR11–756, table 3)
- F<sub>e</sub>dt** **Microtonalite (late Oligocene?)**—Gray, fine-grained, massive, poorly exposed rock consisting of a mosaic of about 20 percent anhedral quartz, 65 percent very sericitized plagioclase, 8 percent biotite (almost entirely altered to chlorite, epidote, and hematite), 4 percent opaque minerals, and 3 percent secondary muscovite. Occurs only as float, just west of hypabyssal rhyolite pluton (F<sub>e</sub>rc) near southeast corner of the map area
- F<sub>e</sub>wf** **Felsic plutons and dikes of Winfield Peak and Middle Mountain of Fridrich and others (1998) (middle Eocene)**—White to pale-gray, dull to vitreous, mostly porphyritic, very fine-grained to aphanitic, massive to weakly flow-foliated dikes and small stocks mainly south of Twin Lakes Reservoir. All samples contains about 2 percent equant quartz phenocrysts as long as 2 mm, and most samples contain 2–3 percent white, strongly sericitized plagioclase phenocrysts as long as 2 mm. Locally, vuggy or vesicular. Some dikes have non-porphyritic chilled margin. In places weathers pinkish-orange. Rhyodacitic or dacitic in composition. New <sup>40</sup>Ar/<sup>39</sup>Ar age on sanidine(?) is 42.25±0.13 Ma (sample GR11–725, table 3)
- F<sub>e</sub>Kr** **Rhyolite dikes (Paleocene or Late Cretaceous)**—Two white to light-gray, glassy to finely crystalline rhyolite dikes as wide as about 10 m containing phenocrysts of bipyramidal quartz as long as about 5 mm, feldspar, and biotite; fracture is conchoidal and glassy. Dikes are mapped about 1 km northwest and about 2 km north of the town of Granite. Whole-rock K-Ar age for dike east of Arkansas River (table 3) is 65.3±2.4 Ma (sample USGS(D)G–A–82, Marvin and others, 1989)
- F<sub>e</sub>Ki** **Felsic porphyry intrusive (Paleocene or Late Cretaceous)**—One small stock east of Arkansas River in SE¼, sec. 19, T. 11 S., R. 79 W. near the eastern boundary of the map area. Consists of very pale gray, massive, porphyritic felsite composed of about 65 percent very fine-grained sugary white matrix, about 20 percent anhedral glassy quartz phenocrysts as long as 3 mm, 10 percent muscovite flakes as long as 3 mm, and 5 percent chalky-white (sericitic) plagioclase as long as 2 mm; no visible mafic minerals. Also contains numerous small irregular vesicles, suggesting shallow emplacement. Rock is a high-aluminum rhyolite or dacite porphyry. Undated, but similarity and proximity to dated rhyolite dike (F<sub>e</sub>Kr) suggest Paleocene or Late Cretaceous age

## Proterozoic Igneous and Metamorphic Rocks

- YXp Pegmatite (Mesoproterozoic and (or) Paleoproterozoic)**—Coarse-grained to very coarse grained, white to light-pink, inequigranular quartz-feldspar-mica rock that forms irregularly shaped, commonly zoned dikes and other intrusive bodies that cut all Proterozoic rocks in map area, but predominantly cuts Mesoproterozoic rocks. Small unmapped bodies cut rocks as old as Paleoproterozoic. Some microcline crystals may be longer than 0.5 m; mica is mostly biotite, but locally includes or is composed entirely of muscovite. Accessory minerals include tourmaline, garnet, and opaque minerals. Unit consists predominantly of pegmatite, which commonly grades into and is intimately mixed with aplite, which locally forms separate dikes and other intrusive bodies. Aplite is similar in composition to pegmatite but is a pinkish-tan, fine- to medium-grained, leucocratic, equigranular rock
- Ygr Monzogranite and quartz monzonite (Mesoproterozoic)**—Small stocks of massive to foliated, generally gray to pinkish-gray, light-brown weathering, medium-grained monzogranite and subordinate quartz monzonite. Intrudes Mesoproterozoic rocks (predominantly granite of Granite, unit Ygrg) so age is Mesoproterozoic; may be satellite bodies of monzogranite of Langhoff Gulch (Ygl)
- Ygl Monzogranite of Langhoff Gulch (Mesoproterozoic)**—Gray, medium-grained, equigranular to inequigranular, massive to weakly flow-foliated, syenogranite, monzogranite, and granodiorite. Contains about 25–35 percent quartz, 20–40 percent plagioclase, 12–45 percent microcline, 4–20 percent biotite, 1–2 percent muscovite, trace to 1 percent sphene, and traces of zircon and apatite; locally contains sparse rutile and garnet. Modal analyses from 19 stained slabs indicate predominantly monzogranite composition. U-Pb zircon age from location, about 5 km southeast of the map area, is 1,445±7 Ma (Kellogg and others, 2013). In most locations unit weathers to distinctive reddish-brown to light grayish-brown, in rounded to blocky outcrops
- Yglf Strongly foliated phase**—Highly strained monzogranite of Langhoff Gulch that intruded, and appears to be interlayered with, inliers of biotite gneiss (Xb). Rock is gray, medium-grained, strongly foliated and locally contains zones of augen gneiss. Mapped in one area just east of Arkansas River, northeast of Lake Creek
- Ygrg Granite of Granite (Mesoproterozoic)**—Medium- to dark-gray to grayish-pink, coarse-grained, massive to moderately foliated monzogranite in the southern and eastern parts of the map area; locally porphyritic, with potassium feldspar phenocrysts as long as about 3 cm. Tends to be more mafic to west, where rock grades into granodiorite mapped as Paleoproterozoic Denny Creek Granodiorite (Fridrich and others, 1998) that was originally named Denny Creek Granodiorite Gneiss (Barker and Brock, 1965). U-Pb zircon (SHRIMP) age from an area north of Clear Creek that was previously mapped as Denny Creek Granodiorite is 1,434±7 Ma (sample GR10–511, table 3), statistically identical to an age from a monzogranite sample near the town of Granite of 1,436±6 Ma (sample WRP–03–6B, table 3). These two ages are statistically identical to those from the granite of Elephant Rock (Kellogg and others, 2013), mapped 6.5 km southeast of the map area which in addition to the similarity in petrography, suggests that both plutons were derived from the same magma source. Granite of Granite is mapped extensively south of the Twin Lakes Reservoir and in one small outcrop along the north shore of the reservoir
- Xgg Granitic gneiss (Paleoproterozoic)**—Light- to medium-gray, fine- to medium-grained, weakly to strongly foliated monzogranite, granodiorite, tonalite, and trondhjemitic gneiss. Composed mostly of 25–30 percent undulatory quartz, 35–60 percent oligoclase, trace to 25 percent microcline, 5–15 percent biotite, 1–2 percent opaque minerals, 0–3 percent hornblende, 0–1 percent muscovite, and trace amounts of zircon, apatite, and garnet. Granitic gneiss is closely interlayered with biotite gneiss (Xb) and leucocratic gneiss (Xlg). Grades in places to layered gneiss resembling biotite gneiss (Xb). Similar rocks in the Gore and Tenmile Ranges, about 50 km north of the map area, (Kellogg and others, 2011) and in the Front Range, about 90 km northeast of the map area, (Kellogg and others, 2008) have SHRIMP U-Pb zircon ages between about 1,770 Ma and 1,755 Ma. Mapped in one 5-km long, northeast-trending body about 1 km south of the Twin Lakes Reservoir

- Xlg **Leucogranite gneiss (Paleoproterozoic)**—Light-gray to almost white, medium-grained, moderately foliated monzogranite or granodiorite orthogneiss, with less than 5 percent biotite. Foliation strikes east-northeast. Mapped east and west of Flume Creek, about 1 km south of Twin Lakes Reservoir
- Xb **Biotite gneiss (Paleoproterozoic)**—Mostly gray to dark-gray, fine- to medium-grained, moderately foliated to strongly foliated and layered biotite gneiss. Migmatitic in most places; contains numerous light-colored layers and lenses (leucosomes) typically 0.1–10 cm thick, although locally may be much thicker. In most outcrops, leucosome bodies form less than half of rock, have sharp contacts with host rock, show much pinch and swell, and are locally strongly folded. Leucosomes are composed of equigranular, massive to weakly foliated, white to light-gray microcline-plagioclase-quartz rock containing less than 5 percent biotite; accessory minerals are muscovite, opaque minerals, sphene, apatite, garnet, and zircon. Forms large inliers in monzogranite of Langhoff Gulch (Ygl) along and east of the Arkansas River and a large northeast-trending body about 1.5 km south of the Twin Lakes Reservoir. Some areas mapped as biotite gneiss include small bodies of granitic rock. Unit undated in map area, but similar rocks in the Gore Range and Front Range north and northeast of the map area, respectively, have U-Pb melanosome ages on zircons of about 1,785–1,750 Ma (Kellogg and others, 2008, 2011). These ages represent those of the source rocks from which detritus that composes the sedimentary protolith was derived. In these same areas, U-Pb zircon geochronology indicates that formation of migmatite (partial melting) occurred about 1,695 Ma (Kellogg and others, 2011)
- Xhg **Hornblende gneiss and amphibolite (Paleoproterozoic)**—Dark-gray to black, weakly to strongly foliated, layered, mostly medium grained, hornblende-plagioclase gneiss and amphibolite, locally containing variable, but small, amounts of biotite, quartz, and augite. Commonly has black-and-white mottled texture due to weathered plagioclase (white) and hornblende (black). Amphibolite contains >50 percent amphibole. Closely interlayered in places with more felsic gneissic rocks. Occurs as one small body about 5 km south of the Twin Lakes Reservoir. Many small, unmapped bodies of Xhg are included with other units, such as biotite gneiss (Xb)
- Xgb **Gabbro (Paleoproterozoic)**—Very dark green hornblende-biotite gabbro containing about 50 percent medium-grained calcic plagioclase, hornblende, and biotite in a dark-green, chert-like (serpentinized?) matrix. Exposed as one small (about 140 m long) intrusive body, interpreted as injected along or near contact between granite of Granite (Ygrg) and monzogranite (Ygr) in southwest part of the map area

16 Geologic Map of the Granite 7.5' Quadrangle, Lake and Chaffee Counties, Colorado



**Figure 2.** Geographic setting of the Granite 7.5' quadrangle. Quadrangle boundary shown by yellow rectangle. Channel of Arkansas River and channel of East Fork of Arkansas River are approximately located.

**Table 2.** Approximate height, in meters, of the top of gravelly stream alluvium of middle Pleistocene to Holocene age above the Arkansas River and its major east-flowing tributaries near the Arkansas River in the Granite 7.5' quadrangle.

[—, no deposits along stream]

Map unit	Arkansas River	Lake Creek	Clear Creek	Cache Creek	Corske Creek
Qva	1–2	1–2	1–2	1–2	1–2
Qopy	2–12	2–7	2–3	2–3	2–3
Qopo	18–24	18	—	—	—
Qoby	37	—	—	—	—
Qobo	43	—	—	—	—
Qoey	55–61	—	—	—	—
	67–72	—	—	—	—
	79–87	—	—	—	—
Qoeo	95–110	—	—	—	—

**Table 3.** Summary of isotopic ages for bedrock in the Granite 7.5' quadrangle.

[All  $^{40}\text{Ar}/^{39}\text{Ar}$  analyses performed in Denver, Colorado, by M.A. Cosca (USGS); all  $^{207}\text{Pb}/^{206}\text{Pb}$  analyses performed in joint USGS-Stanford SHIRMP facilities in Menlo Park, California, by W.R. Premo (USGS); whole-rock K-Ar age by Marvin and others (1989). Abbreviations: feld, feldspar; hb, hornblende; san, sanidine; WR, whole rock; zr, zircon]

Sample No	Rock name	Map symbol	Location (UTM)*	Mineral analyzed	Technique	Age Ma
GR11–711 <sup>1</sup>	vesicular andesite flow	Feva	0385593E 4324406N	hb	$^{40}\text{Ar}/^{39}\text{Ar}$	24.12±0.12
GR11–756 <sup>1</sup>	hypabyssal rhyolite	Ferc	0390236E 4318028N	feld (san?)	$^{40}\text{Ar}/^{39}\text{Ar}$	24.3± 0.6
GR11–725 <sup>1</sup>	rhyolite or dacite porphyry	Fewf	0382206E 4324141N	feld (san?)	$^{40}\text{Ar}/^{39}\text{Ar}$	42.25±0.13
USGS(D)G–A–82 <sup>2</sup>	rhyolite dike	Fekr	0391371E 4324395N	WR	K-Ar	65.3±2.4
GR10–511 <sup>1</sup>	granite of Granite	Ygrg	0385707E 4319201N	zr	$^{207}\text{Pb}/^{206}\text{Pb}$	1,434±7
WRP–03–6B <sup>3</sup>	granite of Granite	Ygrg	0390768E 4321487N	zr	$^{207}\text{Pb}/^{206}\text{Pb}$	1,436±6

<sup>1</sup>K.S. Kellogg sample number<sup>2</sup>Marvin and others (1989) sample number<sup>3</sup>W.R. Premo sample number

\*Datum: North America 1927

## Geologic Setting of the Granite 7.5' Quadrangle

The Granite quadrangle spans the upper Arkansas River valley, which is structurally controlled by normal faults within and along a north-northwest-trending graben that is part of the northern extent of the Rio Grande rift (Tweto, 1979a). The Rio Grande rift system is a zone of crustal extension that extends from northern Mexico (Chapin, 1979) to southern Wyoming (Mears, 1998), and possibly as far north as central Wyoming (Flanagan and Montagne, 1993). Paleoproterozoic and Mesoproterozoic basement rocks are exposed on both sides of the Arkansas River valley. Marine sediments, and mafic and felsic volcanic rocks were deposited locally in the central Rocky Mountain region, including the map area, between about 1,780 and 1,740 Ma. At about 1,715 to 1,700 Ma (Premo and others, 2007), these sediments and rocks were metamorphosed to amphibolite grade and were intruded and deformed by mostly calc-alkalic granitic rocks. These layered metasedimentary and metaigneous rocks commonly include biotite gneiss, granitic gneiss, hornblende gneiss, and minor amphibolite. Zircons from biotite gneiss of the Gore Range, about 50 km north of the map area, are 1,785–1,750 Ma (Kellogg and others, 2011). These are the ages of the source rocks from which the zircons were derived and therefore represent the maximum age of deposition of the sediments. Following metamorphism at about 1,750–1,730 Ma during the Paleoproterozoic, batholiths and smaller bodies of mostly granodiorite and monzogranite, referred to as the Routt Plutonic Suite of Tweto (Tweto, 1978), intruded the older layered rocks in and near the Granite quadrangle. By far the most voluminous intrusive rocks in the quadrangle are granitic rocks of the Berthoud Plutonic Suite of Tweto (Tweto, 1978), which were emplaced during the Mesoproterozoic. These rocks include the granite of Granite (Ygrg, two U-Pb zircon ages of 1,436±6 Ma and 1,434±7 Ma, table 3 and the monzogranite of Langhoff Gulch (Ygl, one U-Pb zircon age of 1,446±50 Ma; Kellogg and others, 2013).

No pre-Neogene sedimentary rocks are exposed in the Granite quadrangle (Tweto and Reed, 1973), probably due to uplift and erosion of the Sawatch Range during the Laramide orogeny or during post-Laramide regional epeirogenic uplift (Eaton, 2008). A thick section of Paleozoic rocks cut chiefly by normal faults is preserved on the east flank of the Mosquito Range immediately northeast of the Granite quadrangle (Tweto and others, 1978).

Numerous dikes and small intrusive stocks intrude the Proterozoic rocks in the quadrangle. These dikes may be as old as Late Cretaceous, based on one whole rock K-Ar age of 65.3±2.4 Ma (table 3), from a rhyolite dike 2 km north of the town of Granite (Marvin and others, 1989). Dikes associated with the Twin Lakes pluton of Paleocene age have not been identified in the Granite quadrangle. The Twin Lakes pluton, emplaced just west of the Granite quadrangle, is composed chiefly of strongly porphyritic biotite granodiorite and associated compositionally layered rocks and aplite bodies

(Fridrich and others, 1998). One of the diorite dikes related to the pluton, about 15 km west of the Granite quadrangle, has a  $^{40}\text{Ar}/^{39}\text{Ar}$  weighted-mean age on hornblende of 63.8±1.4 Ma (Fridrich and others, 1998). A cluster of felsic dikes and small plutons of middle Eocene age were emplaced mainly south of the Twin Lakes Reservoir. One of the felsic plutons (Fwf) has a  $^{40}\text{Ar}/^{39}\text{Ar}$  age of 42.25±0.13 Ma. A small hypabyssal rhyolite pluton (Frc), with a  $^{40}\text{Ar}/^{39}\text{Ar}$  age of 24.3±0.6 Ma, is exposed near the southeast corner of the quadrangle. A small andesite lava flow (Fva), with a  $^{40}\text{Ar}/^{39}\text{Ar}$  age of 24.12±0.12 Ma, overlies Paleoproterozoic rock 1 km south of the Twin Lakes Reservoir (see table 3).

One small felsic stock (FKi) and two rhyolite dikes (FKr) north of the town of Granite were emplaced during the Laramide orogeny, a period of contractional tectonism and mountain building, that began in the area of the present-day Sawatch Range about 72–70 Ma during the Late Cretaceous (Tweto, 1975, 1980). Laramide structural features, such as thrust and reverse faults, have not been documented in the Granite quadrangle, but are regionally common (for example, Tweto, 1975, 1980; Kellogg and others, 2004).

Dominantly north-northwest-trending normal faults, that are down-to-the-east in the western two-thirds of the Granite quadrangle and down-to-the-west in the eastern one-third, resulted from crustal extension that probably began during the late Oligocene. Normal faults are characteristic of the northern part of the Rio Grande rift. The oldest syn-rift volcanic rocks in southern Colorado and northern New Mexico are about 26 Ma (Lipman and Mehnert, 1975; Chapin, 1988; Thompson and others, 1991; Chapin and Cather, 1994); they mark the approximate time of initial rifting. Three fission-track ages on apatite (20.0 ± 2.9, 19.9 ± 4.4, and 16.0 ± 2.4 Ma) for rocks of the Twin Lakes pluton, collected about 2–6 km west of the Granite quadrangle, are cooling ages that reflect the development of the upper Arkansas valley graben and associated significant uplift, erosion, and westward tilting of the east margin of the Sawatch Range during Miocene time (Bryant and Naeser, 1980). The age of the extensional faulting in the Granite quadrangle is not well constrained, and most fault scarps are highly degraded. Lidar imagery suggests that two, short, inferred normal faults about 1 km west of the Arkansas River may displace deposits as young as middle Pleistocene. Evidence of younger faulting is well expressed along the eastern margin of the Sawatch Range, near Cottonwood Creek and Eddy Creek, about 20 and 33 km south of the Granite quadrangle, respectively, where neotectonic studies suggest at least six surface-faulting events during the past 150,000–100,000 years; a recurrence interval of 10,000–40,000 years for significant surface displacement, and youngest displacement during the Holocene (Ostenaar and others, 1980, 1981).

The Arkansas valley within the Granite quadrangle is locally underlain by at least 240 m of basin-fill sediments of the Dry Union Formation (Nd) of Miocene and lower Pliocene(?) age. These deposits consist of unconsolidated to moderately consolidated fluvial and fan deposits composed chiefly of sand and gravel deposited during a period of

widespread extension-related basin development and coeval basin sedimentation that is contemporaneous, in part, with sediments of the Santa Fe Group, Troublesome Formation, and Browns Park Formation in northern New Mexico and northern Colorado (Izett, 1975; Steven, 2002; McMillan and others, 2006).

Surficial deposits of Pleistocene and Holocene age are widespread in the Granite quadrangle, particularly in the valleys of Lake Creek and Clear Creek and on slopes underlain by the Dry Union Formation. During the Pinedale, Bull Lake, and pre-Bull Lake glaciations, large glaciers headed near the crest of the Sawatch Range west of the Granite quadrangle. They flowed 31 km down the valley of Lake Creek and 23 km down the valley of Clear Creek and deposited abundant till (Qt<sub>py</sub>, Qt<sub>po</sub>, Qt<sub>b</sub>, and Qt<sub>e</sub>) and outwash gravel (Q<sub>opy</sub>, Q<sub>opo</sub>, Q<sub>oby</sub>, Q<sub>obo</sub>, Q<sub>oey</sub>, and Q<sub>o eo</sub>) that formed large moraines and extensive outwash terraces in and near the upper Arkansas River valley (Nelson and Shroba, 1998).

Outwash deposits formed during the Pinedale, Bull Lake, and pre-Bull Lake glaciations in the Granite quadrangle were deposited by the Arkansas River and by eastward-flowing tributary streams. The lithologic composition of the clasts in these deposits primarily reflects that of the main bedrock units in the glaciated headwaters of these streams. Outwash deposited by the Arkansas River is rich in clasts of granitic composition, porphyry, and pegmatite. It also contains subordinate amounts of light-colored, aphanitic intrusive rocks and gneiss, and a minor amount of quartzite. Outwash deposited by eastward flowing tributary streams are rich in clasts of granitic composition (including the strongly porphyritic biotite granodiorite of the Twin Lakes pluton) and gneiss. Deposits along Lake Creek and Corske Creek also contain a minor amount of ash-flow tuff, rhyolite tuff, and other volcanic rocks of the Grizzly Peak caldera (Fridrich and others, 1991). Deposits along Clear Creek and Cache Creek also contain a minor amount of metasedimentary gneiss, schist, and thinly bedded gray quartzite (Fridrich and others, 1998). Clasts in outwash deposited by the Arkansas River are commonly angular to subrounded; whereas, those in outwash deposited by eastward-flowing tributary streams are commonly angular and subangular.

The glacier in the valley of Clear Creek impounded the Arkansas River and created large lakes in the Arkansas valley upstream of the glacier twice during the Pinedale glaciation, at about 19 and 18 ka (Young and others, 2011). However, these lakes must have been short lived because neither the deposits that accumulated in these large lakes nor any geomorphic features (such as shorelines) associated with these lakes were observed within the Granite quadrangle. Floods caused by the failures of glacial-ice dams in the valley of Clear Creek at its confluence with the Arkansas River transported boulders greater than 10 m in mean diameter (Lee, 2010), and had estimated peak discharges of between 21,000 and 65,000 m<sup>3</sup>/s (Brugger and others, 2011).

Deposits of sandy silt and silty sand at least 20 m thick, penetrated by drill holes beneath deposits of till and outwash along the axis of the Twin Lakes dam (Tweto, 1978;

D.A. Ostenaar, S.L. Losh, and A.R. Nelson, Bureau of Reclamation, Denver, Colo., written commun., 1981), probably record a proglacial or ice-marginal lake or lakes impounded by the glacier or by till in the valley of Lake Creek during the Pinedale glaciation. The only exposed glacio-lacustrine deposit, too small to show at map scale, is a gray massive silt exposed on the north shore of the Twin Lakes Reservoir about 2 km northwest of the Twin Lakes dam. This deposit was previously considered to be type-O Pearlette volcanic ash (Tweto, 1987); however, it only contains scattered shards of volcanic glass (J.L. Slate, USGS, oral commun., 2011) that probably were derived from volcanic ash in the Dry Union Formation.

Mass-movement deposits transported by creep, debris flows, landslides, rockfall, and solifluction are locally common in the Granite quadrangle. Large landslide deposits (Q<sub>ls</sub>) along and north of the Twin Lakes Reservoir as well as the large landslide on the north side of Cache Creek are probably composed chiefly of material derived from sediments of the Dry Union Formation. The two large landslide deposits near the southwest corner of the Granite quadrangle are composed of rock fragments derived from granite bedrock that may have been weakened by faulting and deep-seated gravitational spreading.

Large placer-tailings deposits (pt) in the Granite quadrangle were formed between about 1860 and 1950 by hydraulic mining and other mining methods used to concentrate native gold (Parker, 1974a, 1974b). This gold that was produced primarily from glacial outwash and fluvial gravel in the Dry Union Formation, probably was derived chiefly from gold-bearing veins in bedrock of the Sawatch Range. The largest placer-tailings deposits are along Cache Creek, west of the Granite Cemetery. Smaller deposits are present along Cache Creek east of the Granite Cemetery, along Lake Creek near its confluence with the Arkansas River, along the low reach of Corske Creek, and along minor, unnamed washes near the mouth of Lake Creek and in the area between Lake Creek and Cache Creek.

## Lithologic Characteristics of the Dry Union Formation in and near the Granite 7.5' Quadrangle

Sediments of the Dry Union Formation (Nd) record the depositional history and architecture of extension-related structures, including the Blue River half graben near Breckenridge, Colo. (Tweto and others, 1978; Kellogg and others, 2011), the upper Arkansas valley graben between Leadville and Salida, Colo. (Scott, 1975; Scott and others, 1975; Tweto and others, 1978; Kellogg, 1999), the Pleasant Valley graben near Howard, Colo. (Taylor and others, 1975), and the north end of the San Luis basin near Poncha Pass (Van Alstine, 1968; Tweto, 1979b). Tweto (1961) defined the Dry Union Formation over 50 years ago, although much of what is known about the lithology of the Dry Union in the northern

part of upper Arkansas valley graben only comes from brief descriptions by Tweto (1961, 1974c, 1978; Tweto and Case, 1972). The following text summarizes information on the genesis, particle-size characteristics, thickness, clast composition, and color of deposits that compose the upper part of the Dry Union Formation within an area south of Dry Union Gulch (see the USGS Leadville South and Granite 7.5' quadrangles) where sediments of the Dry Union Formation are well exposed (Tweto, 1961; McCalpin and others, 2012; this report).

Fluvial deposits commonly consist of cobble gravel, cobbly pebble gravel, pebble gravel, and sandy pebble gravel with a matrix of coarse sand that locally contains abundant granules. They also include deposits of non-pebbly to pebbly coarse to very coarse sand and finer-grained sediments that contain varying amounts of silt and clay (table 4). A 30-m thick section composed of coarse cobble gravel north of Balltown contains a few thin (30 cm) lenses of coarse and very coarse sand and a few rounded boulders about 30–50 cm in diameter. Deposits of fluvial gravel just west of the Granite Cemetery locally contain thin (25–70 cm), roughly horizontal zones of light-orange (iron?) and zones of black (manganese?) coatings on clasts. The formation of these orange and black zones may be due to oxidation and reduction reactions associated with former fluctuating water tables. Gravelly fluvial deposits north of Balltown are rich in granitic clasts, whereas those near the Granite Cemetery also contain subordinate amounts of light-colored aphanitic intrusive rocks and a minor amount of white quartzite. White quartzite clasts, which com-

monly have thick, tan oxidation rinds, probably were derived from sources north of the map area east of the Arkansas River (Tweto, 1974a, 1974b). Clasts in fluvial deposits are angular to subrounded and typically have a light-yellow patina not observed on clasts in deposits of outwash of late and middle Pleistocene age within the map area.

Stream-channel alluvium deposited by ephemeral and intermittent streams are similar in particle size to fluvial deposits, although the former commonly lacks deposits of cobble gravel (table 4) and subrounded clasts, and commonly contains more deposits of pebbly sand and sand than do typical fluvial deposits. Deposits of gravelly alluvium west of the Granite Cemetery are rich in light-colored aphanitic intrusive igneous rocks, clasts of granitic composition, vein quartz, and rhyolite tuff or shallow intrusive rock. Gravelly alluvium in other parts of the map area are rich in granitic clasts.

Deposits of sheetflood alluvium commonly consist of fining upward, rhythmically bedded couplets composed of alternating layers of mud and either pebble gravel or pebbly coarse sand. Muddy upper layers commonly consist of slightly clayey, very silty, very fine and fine sand or very fine sandy silt (table 4). Pebbly lower layers commonly consist of slightly sandy, granule-rich, pebble gravel with a medium to very coarse sand matrix, and locally consist of pebbly coarse and very coarse sand. The muddy sediments that compose some of the upper layers may have been deposited by hyperconcentrated flows. Muddy layers locally show evidence of minor small-scale syndepositional soft-sediment deformation as

**Table 4.** Estimated particle sizes of unconsolidated to moderately consolidated sediments of the Dry Union Formation in the Granite 7.5' quadrangle.

[Abbreviations: FD, fluvial deposits; SA, stream-channel alluvium deposited by ephemeral and intermittent streams; SF, sheetflood alluvium; DF, debris-flow deposits; Sl, slightly; X, observed at one or more sites —, not observed]

Sediments	FD	SA	SF	DF
Cobble gravel	X	—	—	—
Cobbly pebble gravel	X	X	—	—
Cobbly sandy pebble gravel	—	—	—	X
Pebble gravel	X	X	X	—
Sandy pebble gravel	X	X	X	—
Non-pebbly to pebbly coarse and very coarse sand	X	X	X	—
Pebbly, sl. clayey, silty, coarse and very coarse sand	—	—	—	X
Pebbly, clean to sl. silty, medium to very coarse sand	X	—	—	—
Sl. pebbly, clean to silty, very fine to very coarse sand	—	X	—	—
Sl. silty to very silty, very fine to medium sand	X	—	—	—
Sl. clayey, very silty, very fine and fine sand	—	X	X	—
Very fine sandy silt	X	X	X	—

well as evidence of post-depositional down-slope creep that may have occurred under periglacial conditions during the Pleistocene. Clasts in sheetflood deposits commonly are very angular to angular.

Debris-flow deposits fine upward, commonly with an upper layer composed of pebbly, slightly clayey, silty, locally granule-rich, coarse and very coarse sand, and a lower layer composed of cobbly, sandy pebble gravel (table 4). Clasts in debris-flow deposits commonly are very angular to angular, and many are randomly oriented.

The matrix colors of deposits of the Dry Union Formation tend to vary with particle size (for particle sizes used see table 1). Deposits rich in silt commonly are unoxidized or reduced, probably due in part to their relatively high moisture-holding capacity (table 5). They are light gray (5Y 7/2 and 2.5Y 7/2) and pale yellow (5Y 7/3), and commonly have small yellow (10YR 8/6) and reddish-yellow (7.5YR 6/8) mottles. Deposits of sand and gravel are oxidized, typically very pale brown (10YR 7/3), pale brown (10YR 6/3), very pale brown (10YR 7/4), and light-yellowish brown (10YR 6/4) probably due in part to their relatively low moisture-holding capacity and the weathering of iron-bearing minerals such as biotite. Some deposits of sand and gravel are more strongly oxidized, and are pink (7.5YR 7/4), light brown (7.5YR 6/4), and strong brown (7.5YR 5/6).

Although thin beds and lenses of volcanic ash are locally exposed in the Dry Union Formation in the western portion of the Leadville mining district (Tweto, 1961) and near Salida, Colo. (Van Alstine, 1969, 1974; Shannon and McCaLpin, 2006), thin lenses of volcanic ash or fine ashy sediment were only observed within the map area by Shroba in 1974 in a large trench excavated for penstocks near the southeast corner of the Mount Elbert Forebay, now covered by artificial fill (af). A few lenses of volcanic ash were penetrated

by drill holes near the north end of the trench (D.A. Ostenaar, S.L. Losh, and A.R. Nelson, Bureau of Reclamation, Denver, Colo., written commun., 1981).

Eolian, lacustrine, or playa deposits were not observed within the map area. However, deposits of sandy silt and clay were penetrated by drill holes beneath and slightly north of the Mt. Elbert power plant on the north shore of the Twin Lakes Reservoir (D.A. Ostenaar, S.L. Losh, and A.R. Nelson, Bureau of Reclamation, Denver, Colo., written commun., 1981). Silty sediments (Nd) at least 11.5 m thick, which contain one or more thin (35 cm) lenses of volcanic ash and one or more thin (40 cm) lenses of granule-rich very coarse sand, are locally exposed on the north side of Box Creek about 1 km north of the map area. These sediments may have been deposited in a body of water or in a wetland environment on a floodplain during shallow overbank flooding in an area away from the active stream channel. They lack the induration and strong reddish brown to red colors characteristic of sediments deposited in alkaline, saline water (Chapin and Lindley, 1986). Some of the deposits beneath and near the power plant and on the north side of Box Creek may be similar in origin to the thick fine-grained deposits, penetrated by miners in deep (25–245 m) mine shafts excavated in the Dry Union Formation in the western part of the Leadville mining district. These fine-grained deposits were originally called “lake beds” (Emmons and others, 1886; Emmons and Irving, 1907; Emmons and others, 1927). They were later described as consisting predominantly of beds of sandy and pebbly silt, lenses of sand, clay, and gravel, and thin beds of volcanic ash deposited chiefly in alluvial fans (Tweto, 1961, 1974c). Some of the silt-sized material in thick fine-grained deposits in the western part of the Leadville mining district may be similar in origin to that of abundant siltstone beds of the upper Oligocene to upper Miocene Troublesome Formation exposed

**Table 5.** Dry Munsell colors of unconsolidated to moderately consolidated sediments of the Dry Union Formation in the Granite 7.5' quadrangle.

[Abbreviations: FD, fluvial deposits; SA, stream-channel alluvium deposited by ephemeral and intermittent streams; SF, sheetflood alluvium; DF, debris-flow deposits; X, dry matrix color observed at one or more sites; —, color not determined or not observed]

Color	Munsell notation	FD	SA	SF	DF
Pale yellow	5Y 7/3	—	X	—	—
Light gray	5Y 7/2	X	X	—	—
Light gray	2.5Y 7/2	—	X	—	—
Very pale brown	10YR 7/4	X	X	—	—
Very pale brown	10YR 7/3	X	X	X	X
Light yellowish-brown	10YR 6/4	—	X	—	—
Pale brown	10YR 6/3	X	—	—	—
Pink	7.5 YR 7/4	X	X	—	X
Light brown	7.5 YR 6/4	—	X	—	—
Strong brown	7.5 YR 5/6	X	—	—	—

in the Fraser and Granby basins and in Middle Park about 90 km north of the map area. Limited particle-size and sorting data for siltstones in the Troublesome Formation in the Fraser basin suggests that these siltstones are composed chiefly of loess or loess-rich alluvium (Shroba and others, 2010).

A few of the deposits composed of very silty, very fine to medium sand south of the Mount Elbert Forebay have blocky, soil-like structure. These deposits may be weakly developed cambic B soil horizons that formed in paleosols. These deposits or incipient soil horizons lack clay films, which are commonly observed in argillic B soil horizons (Soil Survey Staff, 1999). Near Dry Union Gulch about 5 km north of the map area, McCalpin and others (2012) have identified two paleosols with clayey, reddish-brown, argillic B horizons preserved in the exposed upper part of the Dry Union Formation. Neither argillic nor calcic paleosol horizons were observed within the map area.

## Age of the Dry Union Formation

The age of the Dry Union Formation in the Leadville area is not well known (Tweto, 1961). Well exposed beds near Salida, Colo., near the southern end of the upper Arkansas valley graben, locally contain abundant remains of fossil mammals and deposits of volcanic ash that suggest a Miocene age. G.E. Lewis considered the fossils identified near Salida (Van Alstine and Lewis, 1960; Van Alstine, 1969, 1974) such as *Neohipparion* sp. (three-toed horse) to be of Pliocene age, but they are now considered to be of Miocene age based on the current subdivision of Neogene time (U.S. Geological Survey Geological Names Committee, 2010). The oldest fossil mammals collected from the Dry Union Formation near Salida (fig. 2) are of Barstovian age (Steven and others, 1997) and therefore they are probably no older than about 15.9 Ma. The youngest fossil mammal collected from the Dry Union Formation, located in Adobe Park about 5 km northwest of Salida, is *Ceratogaulus hatcheri* (K.C. McKinney, USGS, Denver, Colo., written commun., 2012). The range in age of this extinct rodent is about 6–5 Ma (Hopkins, 2005).

A deposit of gray vitric ash about 1 km south of Salida has a composition that is typical of ashes erupted from the Twin Falls volcanic field in Idaho between about 10.5 and 8.5 Ma (Xu and others, 2001). Two other ash deposits in older, clayey, lacustrine(?) sediments of the Dry Union Formation (Shannon and McCalpin, 2006), about 12 km west of Salida, yielded radiogenic argon ages of  $14.2 \pm 1.4$  and  $13.3 \pm 0.2$  Ma (Hubbard and others, 2001).

The range in age of *Ceratogaulus hatcheri* along with the ages of the volcanic ashes investigated by Xu and others (2001) and Hubbard and others (2001) suggest that some of the sediments in the Dry Union Formation near Salida are coeval with those of the Ash Hollow Formation of the Ogallala Group in the northern part of the High Plains, which ranges in age from 13 to about 6 Ma (Perkins and others, 1995).

The older, clayey ash-bearing sediments west of Salida may be correlative, in part, with the Pojoaque Member of the Tesuque Formation (Santa Fe Group), Troublesome Formation, North Park Formation, and Browns Park Formation. The Pojoaque Member of the Tesuque Formation, at or near its type section north of Santa Fe, New Mexico, contains a tephra layer near the middle of the member that yielded a weighted-mean  $^{40}\text{Ar}/^{39}\text{Ar}$  age of  $13.7 \pm 0.18$  Ma (Izett and Obradovich, 2001). The Troublesome Formation in Middle Park in north-central Colorado contains eight tuffs that have  $^{40}\text{Ar}/^{39}\text{Ar}$  ages on sanidine and plagioclase that range from 23.5 to 11.0 Ma and fossil mammals that range in age from latest Arikareean to early Clarendonian (Izett and Obradovich, 2001). The North Park Formation in North Park in north-central Colorado locally contains a rhyolitic tuff that has a mean  $^{40}\text{Ar}/^{39}\text{Ar}$  age on sanidine of  $28.17 \pm 0.05$  Ma (E.E. Larson, Professor Emeritus, Univ. of Colo. at Boulder, written commun., 2013) and locally contains fossil mammals of late Hemingfordian age (Hail and Lewis, 1960) and Barstovian age (Montagne and Barnes, 1957; Montagne, 1991; Steven and others, 1997). The upper part of the Browns Park Formation, preserved in a half-graben (Mears, 1998) near Saratoga in south-central Wyoming, contains two pumicite layers that have  $^{40}\text{Ar}/^{39}\text{Ar}$  weighted-mean ages of  $15.86 \pm 0.04$  on biotite and  $14.54 \pm 0.04$  Ma on anorthoclase and fossil mammals of Barstovian and Hemingfordian ages (Montagne, 1991). The Browns Park Formation in northwestern Colorado, west of the Park Range, locally contains air-fall tuffs dated by fission-track, K/Ar, and  $^{40}\text{Ar}/^{39}\text{Ar}$  methods that range in age from 24.8 to 7.2 Ma (Izett and others, 1970; Winkler, 1970; Izett, 1975; Naeser and others, 1980; Snyder, 1980a; Snyder, 1980b; Luft, 1985; Honey and Izett, 1988).

The age of the sediments of the Dry Union Formation that underlie the older, clayey, ash-bearing sediments west of Salida is unknown. The former may be correlative, in part, with sediments of the Arikaree Group in western Nebraska that range in age from about 29 to 18 Ma (Swinehart and Diffendal, 1995; Tedford and others, 2004).

The Dry Union Formation near Salida is younger than the Browns Canyon Formation (Van Alstine, 1969). Remnants of the Browns Canyon Formation, composed predominantly of thin-bedded tuffaceous siltstone, are exposed in a 0.2 km<sup>2</sup> area about 13 km northwest of Salida where they contain plant fossils similar to those preserved in sediments of the Creede Formation in southwestern Colorado (Van Alstine, 1969). Sediments of the Creede Formation were deposited between 26.8 and 26.3 Ma ( $^{40}\text{Ar}/^{39}\text{Ar}$  ages on sanidine from underlying and overlying volcanic units, Lanphere, 2000), during the late Oligocene.

The deposition of the Dry Union Formation in the southern part of the upper Arkansas valley graben postdates the inception of rift-related faulting and disruption of a former northeast-flowing stream in the Trout Creek paleovalley (Epis and other, 1976; Tweto, 1978; Leonard and others, 2002), 8 km southeast of Buena Vista, Colo. (fig. 2), and postdates the deposition of the Brown Park Formation (Van Alstine,

1969). Fluvial transport in the paleovalley ceased sometime after clasts composed of rhyolite of Bald Mountain were deposited in sediments of the Wagontongue Formation in South Park (Tweto, 1978), about 20 km northeast of Buena Vista. Transport and deposition of these clasts occurred sometime after the eruption of the rhyolite of Bald Mountain at about 30.4 Ma (on sanidine in older flow) and 28.9 Ma (on anorthoclase in younger flow; McIntosh and Chapin, 2004).

The radiometric ages of the rhyolite of Bald Mountain (McIntosh and Chapin, 2004) suggest that the inception of rift-related faulting and deposition in the area near the Trout Creek paleovalley began after 30.4 or 28.9 Ma. The recognition of similar plant fossils in the Browns Canyon Formation and the Creede Formation suggests that the inception of rift-related faulting and deposition of the Browns Canyon Formation may have begun as early as the late Oligocene. A late Oligocene age for the inception of rift-related faulting and deposition in the upper Arkansas valley graben is in accord with the age of about 26 Ma for the inception of rift-related volcanism recorded by a thick series of mafic lavas preserved in the San Luis Hills in the central part of the San Luis basin (Thompson and others, 1991), about 135 km southeast of the southern margin the upper Arkansas valley graben.

## Brief History of Placer-gold Mining in the Granite 7.5' Quadrangle

Some of the placer-gold mining areas within the map area have produced significant amounts of gold (table 6) and are among the richest in the State of Colorado. The following history was summarized chiefly from reports by Parker (1974a, 1974b, 1992). Gold was first discovered in deposits along Cache Creek in 1859 (Henderson, 1926) or 1860 (Emmons and others, 1886). The news of the placer gold discovery along Cache Creek spread rapidly. By the end of 1860, gold was discovered in deposits elsewhere within the map area, such as near the head of Lost Canyon, in the area between Cache Creek and Lake Creek, and near the mouth of Lake Creek. Between about 1859 or 1860 and 1883, placer gold was produced chiefly by hand methods and by ground sluicing. Gold production along Cache Creek, and possibly at small placer-mining areas north of Cache Creek, was increased in 1863 by water from Lake Creek that was conveyed by the Arlington ditch on the south side of Lake Creek. The Cache Creek ditch, which conveyed water from Big Willis Gulch (about 1 km west of map area) to small placer workings at and east of Flume Creek, probably was completed prior to 1900. Hydraulic mining within the map area began in 1884 with the completion of a ditch, tunnel, and flume, all of which conveyed water from the Clear Creek drainage to the Cache Creek placer-mining area. Hydraulic mining along Cache Creek continued from 1884 to 1917 when mining ceased.

Placer-mining and gold production significantly decreased during and after 1911 as the result of a lawsuit filed by the City of Pueblo and various water districts in the lower Arkansas River valley in 1910 against the company that was carrying out hydraulic mining operations along Cache Creek. The lawsuit, possibly one of the first environmental lawsuits in the State of Colorado, was filed because significant amounts of sediment generated by placer-mining along Cache Creek was flowing into and degrading the water quality of the Arkansas River.

Between 1860 and 1898, gold was produced by hand methods from bouldery outwash deposits of Pinedale age along Lake Creek east of the Twin Lakes. Ditches that conveyed water from Lake Creek facilitated hydraulic mining near the mouth of Lake Creek, and in areas north and south of Lake Creek, between about 1898 and 1908. Abundant boulders in the surficial deposits near the mouth of Lake Creek made hydraulic mining and the extraction of gold difficult.

During the early 1930s and periodically thereafter, gold was produced again by hand methods near the mouth of Lake Creek. A minor amount of dredging was done at the confluence of Lake Creek and the Arkansas River, and along the Arkansas River north of Granite during or slightly after 1927. Although gold may have been discovered along Corske Creek during the 1860s, substantial gold production did not begin until dredging started in 1915. Dredging along and near the lower reach of Corske Creek was carried out during 1915–1917, 1924–1926, and 1947–1950. A dragline was also used to mine gold in this area during 1934–1942.

Placer gold within the map area was probably derived chiefly from gold-bearing veins in bedrock in the Lake Creek drainage (Howell, 1919) and near the head of Lost Canyon, which is located west of the Cache Creek drainage. Gold-bearing veins near Gordon Gulch (Gese and Scott, 1993; Cappa and Bartos, 2007), on the north side of the Lake Creek valley near Parry Peak, may have been a major source of the placer gold that was mined along and near the lower reach of Corske Creek. In placer-gold mining areas underlain by Proterozoic bedrock, such as along Cache Creek, the highest concentrations of gold are located in sediments that are a few centimeters to about 4.5 m above the bedrock. However, in areas underlain by silty and clayey sediments of the Dry Union Formation, such as along Corske Creek, much of the gold is concentrated in outwash deposits about 15 cm or less above the top of the Dry Union Formation.

Minor amounts of gold are still being produced from deposits along Cache Creek and Lake Creek at its confluence with the Arkansas River. Present day recreational prospectors now use gold pans and small portable metal sluice boxes to concentrate and obtain gold from deposits worked about a century ago by miners, who used large heavy wooden sluice boxes and hydraulic mining equipment.

**Table 6.** Summary of mining methods, years of gold production, and estimated minimum total value of gold production in the main areas of placer-gold production in the Granite 7.5' quadrangle. All of the information summarized from Parker (1974a, 1974b, 1992).

[All values for gold produced are those reported at the time of production in U.S. dollars. During the period of gold production, the official U.S. Government price of gold was \$20.67 per troy ounce (\$0.67 per gram) until January 30, 1934, when the price increased to \$35.00 per troy ounce (\$1.13 per gram; Parker, 1974a). Values for the Cache Creek and Lake Creek areas do not include the minor amounts of gold produced in these areas during recent decades.]

Main areas of placer-gold production	Mining methods	Years of gold production	Estimated minimum total values of gold production
head of Lost Canyon	hand methods and ground slucing	1860–1959 <sup>1</sup>	200,000
Cache Creek	hand methods and ground slucing	1860–1883	1,000,000
	hydraulic mining	1884–1917	
Lake Creek	hand methods	1859 or 1860–1867	55,000 <sup>2</sup>
	hydraulic mining	1898–1908	
	hand methods	early 1930s <sup>3</sup>	
Corske Creek	dredging	1915–1917	1,300,000
	dredging	1924–1926	
	dragline	1934–1942	
	dredging	1947–1950	

<sup>1</sup>Gold production between 1859 and 1960 was intermittent. Some gold may have been produced after 1959.

<sup>2</sup>Value as of 1867; value of subsequent gold production is unknown.

<sup>3</sup>Gold has been produced periodically since the early 1930s.

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