

### DESCRIPTION OF MAP UNITS

**SURFICIAL DEPOSITS**

- af Artificial fill (Holocene)**—Unconsolidated soil, sand, and gravel that underlie industrialized floodplain areas of northern Portland, Ore., and Vancouver, Wash.; mounds of sand and minor gravel from channel dredging that flank Columbia River; and earth and crushed rock for highway and railroad beds, levees, and small dams.
- Qf Fan deposits from tributaries (Holocene and Pleistocene)**—Unconsolidated silt, sand, and gravel in small fan-shaped accumulations from steep drainages in Tualatin Mountains. Most fans are younger than 2,000 years, inferred from relation with the Columbia River and Willamette River floodplain deposits (Goww and Qcwf). Poorly exposed, but likely composed of silt, sand, and gravel diameters deposited by debris flow and stratified sediment deposited by streamflow.
- Qa Alluvium of tributary floodplains (Holocene and Pleistocene)**—Unconsolidated sand, gravel, and organo-rich mud along rivers and creeks within Tualatin Mountains. Locally upstream of landslides, indicating deposition in landslide-dammed valley bottoms.
- Qe Eolian deposits (Holocene)**—Unconsolidated, massive fine sand and silt forming dunes and benches that crest as much as 24 m (79 ft) above sea level. Mostly quartzofeldspathic sand that contains muscovite and minor lithic fragments. Forms undulating, apparently wind-fluted topography of southwestern Sauvie Island; locally mantles low benches at base of Tualatin Mountains, where mapped bodies may include the compositionally similar cataclysmic-flood deposits, sand and silt facies (Qfs). No evidence of modern sand accumulation or dune activity exists. Oxidized surface soils are as thick as 2 m on Sauvie Island dunes, indicating stability for hundreds or thousands of years. Well logs indicate that dune sand generally extends to depths of 10 to 20 m (33–67 ft) below sea level, locally to 30 m (100 ft) below sea level, suggesting that they began forming as long as 9,000 years ago according to elevations of dated Holocene floodplain deposits along lower Willamette River (Peterson and others, 2011). Likely formed by easterly winds that entrained Columbia and Willamette River beach and bar sands during periods of low water.
- Columbia River and Willamette River floodplain deposits (Holocene and Pleistocene)**—Unconsolidated sediment of floodplain, islands, and bars of Columbia and Willamette Rivers at elevations mostly less than 10 m (30 ft) above sea level, largely composed of quartz, feldspar, and conspicuous muscovite, indicating Columbia River provenance. Upper elevation of deposits approaches limits of historical flooding before 20th-century river regulation and floodplain diking. Most deposits above modern low-water river level (about 5 ft [1.5 m] above sea level in map area) are younger than 2 ka. Deposits at depth likely to be as old as 16 ka (Baker and others, 2010). Well logs and seismic-reflection profiles show that fine-grained silt fill beneath historic floodplain locally extends to 65 to 70 m (210–230 ft) below sea level in and near map area (Gates, 1994; Pratt and others, 2001; Peterson and others, 2011). These deposits resulted from river aggradation since last-glacial sea-level lowstand of about 16 ka.
- Silt and clay facies—Silt, clay, and minor sand deposited in low-lying areas of Columbia River floodplain at elevations below 3 m (10 ft). Includes areas of drained lake beds on Sauvie Island. Bank exposures, auger holes, and trenches show silt, clay, and organic materials in horizontal laminae typically 1 to 20 mm thick.**
- Sand facies—Sand and silt of ridge and swale topography and natural levees flanking floodplain channel and silt of ridge-and-swale topography and natural levees that flank floodplain channels, commonly as much as 9 m (30 ft) above sea level. Ridges and swales consist of an ebbelon point bars or natural levees, indicating migration of Columbia River and secondary channels like Multnomah Channel. Exposed stratigraphy outside map area shows stratified fine to medium sand, in layers as thick as 30 cm, alternating with silty sand. Subhorizontal and inclined beds indicate deposition by both vertical and lateral accretion. Higher sandy deposits that flank meandering floodplain channels are natural levees and crevasse splays from overbank flooding.**
- Landslide deposits (Holocene and Pleistocene)**—Diameters of angular bedrock and surficial debris transported downslope en masse by semicoherent slumps, rockslides, earthflows, and debris flows. Many mapped landslides head at arcuate scars and exhibit subhorizontal tops, bulbous toes, and hummocky surfaces. Large landslide complexes are present along McCarthy Creek.
- Terrace deposits (Holocene and Pleistocene)**—Unconsolidated silt, sand, and gravel that form benches along McCarthy Creek in Tualatin Mountains. Not exposed at surface but mapped by morphology. Likely consists of stratified deposits of streams and diamicts of debris flows.
- Loess (Holocene? and Pleistocene)**—Massive unconsolidated deposits of light-gray to buff, micaceous, quartzofeldspathic eolian clay, silt, and fine sand on upland areas of Tualatin Mountains; contains granules and small pebbles, locally capped by strongly developed red soils. Forms mantle as much as 12 m thick along ridges southwest of Multnomah Channel that overlies the Sentinel Bluffs Member of the Grand Ronde Basalt (Tgwb). Probably deposited during several episodes throughout late Quaternary time; luminescence ages from similar loess deposits outside map area range from greater than 79 ka to 39 ka. Radiocarbon ages on eolian silt on upland areas near Ridgefield, east of map area, indicate episodic deposition throughout Holocene (Punke and others, 2011). Equivalent to the loess of Trimble (1963) and the Portland Hills Silt of Lentz (1981).

**BEDROCK**

**COLUMBIA RIVER BASALT GROUP**

- Tgwb Sentinel Bluffs Member (Pliocene and/or Miocene)**—Semi-consolidated, well-bedded sandstone, siltstone, claystone, and minor quartzite-bearing conglomerate, pumice-lapilli tuff, and lignite. Planar and trough cross-beds and cut-and-fill structures within sandy horizons indicate fluvial deposition. Age is poorly known but probably is late Pliocene to middle Miocene on basis of relations elsewhere in Portland Basin (Evarts, 2004; Evarts and O'Connor, 2008; Evarts and others, 2009a). Shown in cross section only.
- Tg Grand Ronde Basalt (Miocene)**—Light-gray to black flows of tholeiitic basaltic andesite; vesicular to microvesicular, aphyric to microphyric to very sparsely plagioclase-phryic; exhibit relatively low TiO<sub>2</sub> contents that are characteristic of the Grand Ronde Basalt of the Columbia River Basalt Group (Swanson and others, 1979; Mangan and others, 1979; Beeson and others, 1989; Reidel and others, 1989; Reidel and Tolan, 2013). Flows in this quadrangle are deeply weathered, where exposed to roadcuts and quarries, flows dipshaly (enabular) to blocky to columnar-jointed patterns and vesicular flow logs; pillow lava locally present in part. Typical samples are intergranular to interstitial, containing labile plagioclase, granular clinopyroxene, and Fe-Ti oxide crystals in sparse to abundant dark glass; some flows contain rare plagioclase phenocrysts and glomerocrysts as long as 5 mm. Textures resemble those described by Long and Wood (1986) for correlative flows in Columbia Basin. Basalt issued from vents in eastern Columbia Basin and entered northwestern Oregon and southwestern Washington through wide gap in Cascade Range (Tolan and others, 1989; Wells and others, 1989; Beeson and Tolan, 1990). <sup>40</sup>Ar/<sup>39</sup>Ar age determinations (Barry and others, 2010, 2013; Baksi, 2013) indicate emplacement of Grand Ronde flows at about 16 Ma. Nomenclature is from Reidel (2005) and Reidel and Tolan (2013).
- Tgpc Winter Water Member—Sparingly plagioclase-phryic to glomerophyric basaltic andesite flows that have relatively low MgO contents (3.4–3.9 weight percent) and moderate TiO<sub>2</sub> contents (2.0–2.2 weight percent). Contains scattered plagioclase phenocrysts and glomerocrysts 1 to 3 mm across in aphyric to sparsely microphyric groundmass. Exhibits normal magnetic polarities with low inclinations (<35°). J.T. Hagstrum, written commun., 2010). Stratigraphic position and petrographic, chemical, and paleomagnetic characteristics indicate equivalence to Winter Water Member of Reidel and Tolan (2013), within N2 magnetostatic unit of Swanson and others (1979). At least 30 m thick where exposed in Angell quarry near south edge of map area. Most analyzed samples from map area are highly weathered (denoted by low totals and/or anomalously low iron contents of <11 weight percent FeO\*). Data from fresher samples indicate that two chemically distinguishable flows are present in map area. One has MgO content of about 3.9 weight percent, TiO<sub>2</sub> about 2.05 weight percent, and Ba <600 ppm. Other has MgO content of 3.6 weight percent, TiO<sub>2</sub> >2.10 weight percent, and Ba <600 ppm. Distribution in map area, especially north and west of McCarthy Creek, is extremely uncertain owing to poor exposure, and part of that area may be underlain by the informal Orley member (Tgpc) of Reidel and Tolan (2013) (Madin and others, 2008).**
- Tg Orley member of Reidel and Tolan (2013)—Flows of aphyric, interstitial to intergranular basaltic andesite that has relatively low MgO (3.5–3.6 weight percent) and TiO<sub>2</sub> (1.95 weight percent) contents and normal magnetic polarity (J.T. Hagstrum, written commun., 2010). Stratigraphic position and petrographic, chemical, and paleomagnetic characteristics indicate equivalence to the informal Orley member of Reidel and Tolan (2013), within N2 magnetostatic unit of Swanson and others (1979). Four flow units totaling about 65 m thick are exposed in Angell quarry; base of uppermost flow is pillow lava. Most analyzed samples from map area, even those from quarry, are highly weathered (denoted by low analytical totals and/or iron contents of <11 weight percent FeO\*).**
- Tgpc Grouse Creek member of Reidel and Tolan (2013)—Flows of aphyric basaltic andesite exposed only in Angell quarry where total thickness of three flows is 30 m. Uppermost flow is about 40 m thick and is extremely magnetized; overlies columnar-jointed flow about 6 m thick, which, in turn, overlies about 3 m of hackly jointed lowest flow. Chemical analyses show that only lower two flows are relatively fresh; their compositions resemble those of Grouse Creek flows member mapped elsewhere in western Oregon (R.E. Wells, oral commun., 2014; R.C. Evarts, unpub. data).**

**PALEOGENE SEDIMENTARY ROCKS**

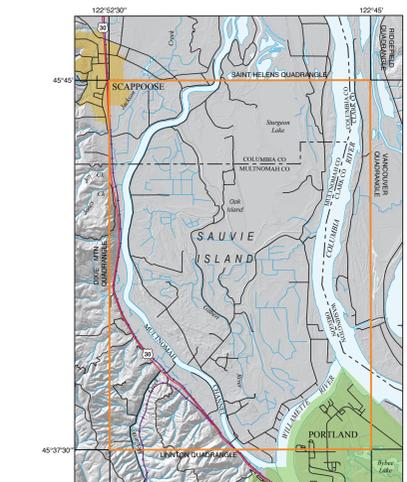
- Tf Scappoose Formation (Oligocene)**—Shallow-marine, tuffaceous, micaceous, arkosic sandstone; underlies the Grand Ronde Basalt to west and northwest of map area (Madin and Niendorf, 2008; M.G. Savaris, oral commun., 2013). Abundant mollusk fauna indicate late Oligocene age (Warren and Nobisrab, 1946). Shown in cross section only.

### EXPLANATION OF MAP SYMBOLS

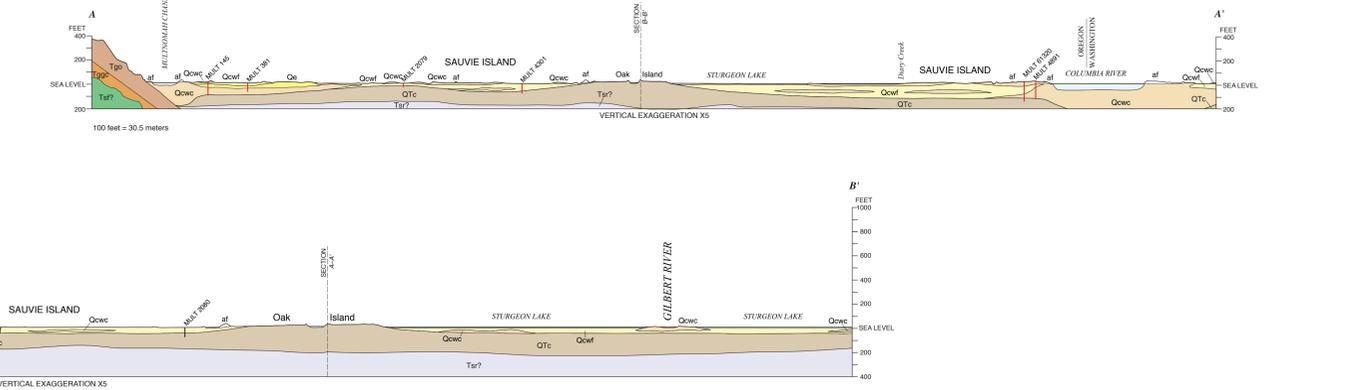
- Contact**—Solid where location is accurate; long-dashed where location is approximate; short-dashed where location is inferred; queried where existence or extent is uncertain.
- Fault**—Long-dashed where location is approximate; short-dashed where location is inferred; dotted where location is concealed; queried where existence or extent is uncertain. Ball and bar on downthrown side. Arrows show relative movement. A, away from observer; T, toward observer (in cross section).
- Anticline**—Trace of axial plane; long-dashed where location is approximate; dotted where location is concealed.
- Syncline**—Trace of axial plane; dotted where location is concealed; queried where existence or extent is uncertain.
- Strike and dip of beds**

### SAMPLE LOCALITIES

- 11** Chemical analysis—Showing map number; see table 1 for analyses.
- 12** Paleomagnetic analysis
- 13** Radiocarbon age—Showing map number; see table 4 for analyses.
- 14** Approximate location of water wells used to construct cross sections—Marked by red vertical lines in cross sections, showing Oregon Department of Water Resources identifier number.



INDEX MAP SHOWING GEOGRAPHIC AND CULTURAL FEATURES OF SAUVIE ISLAND QUADRANGLE (ORANGE OUTLINE) AND VICINITY ON HILLSHADE IMAGE DERIVED FROM LIDAR DATA. ALSO SHOWN ARE NEIGHBORING QUADRANGLES MENTIONED IN TEXT.



## Geologic Map of the Sauvie Island Quadrangle, Multnomah and Columbia Counties, Oregon, and Clark County, Washington

By  
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2016

Base map topography and hydrography compiled from 2009 National Agriculture Imagery Program orthorectified and from lidar survey acquired in 2005 and 2009 by U.S. Army Corps of Engineers and in 2001 by Clark County, Washington.

Cultural data derived from U.S. Census Bureau TIGER files and political boundaries from U.S. Geological Survey 2011 Sauvie Island digital topographic quadrangle.

Data frame projection and 1000-meter grid: Universal Transverse Mercator, Zone 10, North American Datum 1983 (NAD83). Corner ticks for 7.5 quadrangle NADES shown in blue.

Geologic data compiled to NAD27 quadrangle boundaries.

Geology mapped by R.C. Evarts, 2004–2012; J.E. O'Connor, 2006–2012; C.M. Cannon, 2010–2012; and J.P. Madin, 2006; assisted by P.A. Dittman, 2004.

Digital cartography by C.M. Cannon, J.F. Mangano, and K.L. Wheeler.

Edited by Kate Jacques and Taryn A. Linquist.

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