

Structure Contour and Overburden Maps of the Niobrara Interval of the Upper Cretaceous Cody Shale in the Wind River Basin, Wyoming

Pamphlet to accompany
Scientific Investigations Map 3427

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**U.S. Department of the Interior
U.S. Geological Survey**

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Conversion Factors

U.S. customary units to International System of Units

Multiply	By	To obtain
Length		
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
Area		
square mile (mi ²)	2.590	square kilometer (km ²)

Datum

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above sea level.

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By Thomas M. Finn

Introduction

The Wind River Basin is a large northwest-trending structural depression that is nearly 200 miles (mi) long, 70 mi wide, and encompasses about 7,400 square miles (mi²) in central Wyoming (fig. 1). The present-day structure of the Wind River Basin developed during the Laramide orogeny, a period of crustal instability that began during latest Cretaceous time and ended in the early Eocene (Gries, 1983; Love, 1988). Many of the structures are the result of compressional deformation characterized by Precambrian basement-involved thrust faults and strongly folded and faulted anticlines and synclines. The north, west, and south margins of the basin are bounded by basement-cored uplifts referred to as the Washakie, Owl Creek, Bighorn, Wind River, and Granite Mountains uplifts (fig. 2). The Casper arch, a major northwest-trending structural upwarp with a thrust faulted, steeply dipping west limb forms the east margin of the basin (Keefer, 1970) (fig. 2). These uplifts are flanked by highly folded and faulted sedimentary rocks that range from Paleozoic to Paleocene in age, whereas the central part of the basin is covered by nearly flat-lying lower Eocene and undifferentiated Tertiary and Quaternary rocks that mask the structure of the older rocks in the central part of the basin (Keefer, 1966; Keefer, 1970; Love and Christiansen, 1985) (fig. 3).

According to Knight (1897) and Rountree (1984), the first documented account of oil in what is now the State of Wyoming was discovered at a seep along the banks of the Popo Agie River, about 8 mi southeast of the present-day town of Lander near the southwestern margin of the Wind River Basin (fig. 2). For several decades following its discovery in 1833, this seep, referred to as the “great tar spring,” was used by local hunters and fur trappers for medicinal purposes, and later by the military and pioneers for axle grease and lubrication (Knight, 1897; Rountree, 1984). Then in 1884, Mike Murphy completed the first commercial oil well in Wyoming next to the “great tar spring” establishing the first oil production in Wyoming at the Dallas oil field (Biggs and Espach, 1960; Keefer, 1969; Cardinal and others, 1989) (fig. 2). Since then, many important conventional and unconventional oil and gas resources have been discovered and produced from

reservoirs ranging in age from Mississippian through Tertiary (Keefer, 1969; Fox and Dolton, 1989, 1996; De Bruin, 1993; Johnson and others, 1996, 2007) (fig. 2). It has been suggested that various Upper Cretaceous marine shales are the principal hydrocarbon source rocks for many of these accumulations (see, for example: Meissner and others, 1984; Fox and Dolton, 1989, 1996; Johnson and Rice, 1993; Nuccio and others, 1996; and Schelling and Wavrek, 1999, 2001). With recent advances in horizontal drilling and multi-stage hydraulic fracture stimulation, equivalent Upper Cretaceous marine source rock intervals, in particular the Niobrara Formation, have become important continuous (unconventional) shale gas or shale oil objectives in other Rocky Mountain Laramide basins (Matthews, 2011; Sonnenberg, 2011; Williams and Lyle, 2011; Durham, 2012a,b, 2013; Taylor and Sonnenberg, 2014; Hawkins, 2016). In the Wind River Basin, the Niobrara is represented by shales, calcareous shales, marls, siltstones, and sandstones in the lower shaly member of the Upper Cretaceous Cody Shale (Finn, 2017) (fig. 4).

The maps presented in this report were constructed as part of a project carried out by the U.S. Geological Survey (USGS) to better characterize the geologic framework of potential undiscovered continuous (unconventional) oil and gas resources in the Niobrara interval of the Upper Cretaceous Cody Shale in the Wind River Basin in central Wyoming (sheets 1 and 2). The structure contour map is drawn at the base of the “chalk kick” marker bed, a distinctive zone or peak on resistivity logs in the lower 50–300 feet (ft) of the Cody Shale that according to Finn (2017) represents the base of the Niobrara interval in the Wind River Basin (fig. 5). This horizon was selected because it is easily identified on most well logs, is present throughout the basin, and has been identified by the author in about 630 wells (Finn, 2019). Data from an additional 90 wells that penetrated the upper part of the Niobrara interval or overlying Cody Shale were incorporated by projecting thicknesses from nearby wells and estimating a pick for the “chalk kick” marker (Finn, 2019). The structure contouring is based on interpretation of the well data and surface geologic mapping by Keefer (1970), Keefer and Troyer (1964), Love and Christiansen (1985), Green and Drouillard (1994), Sutherland and Hausel (2003), and Johnson

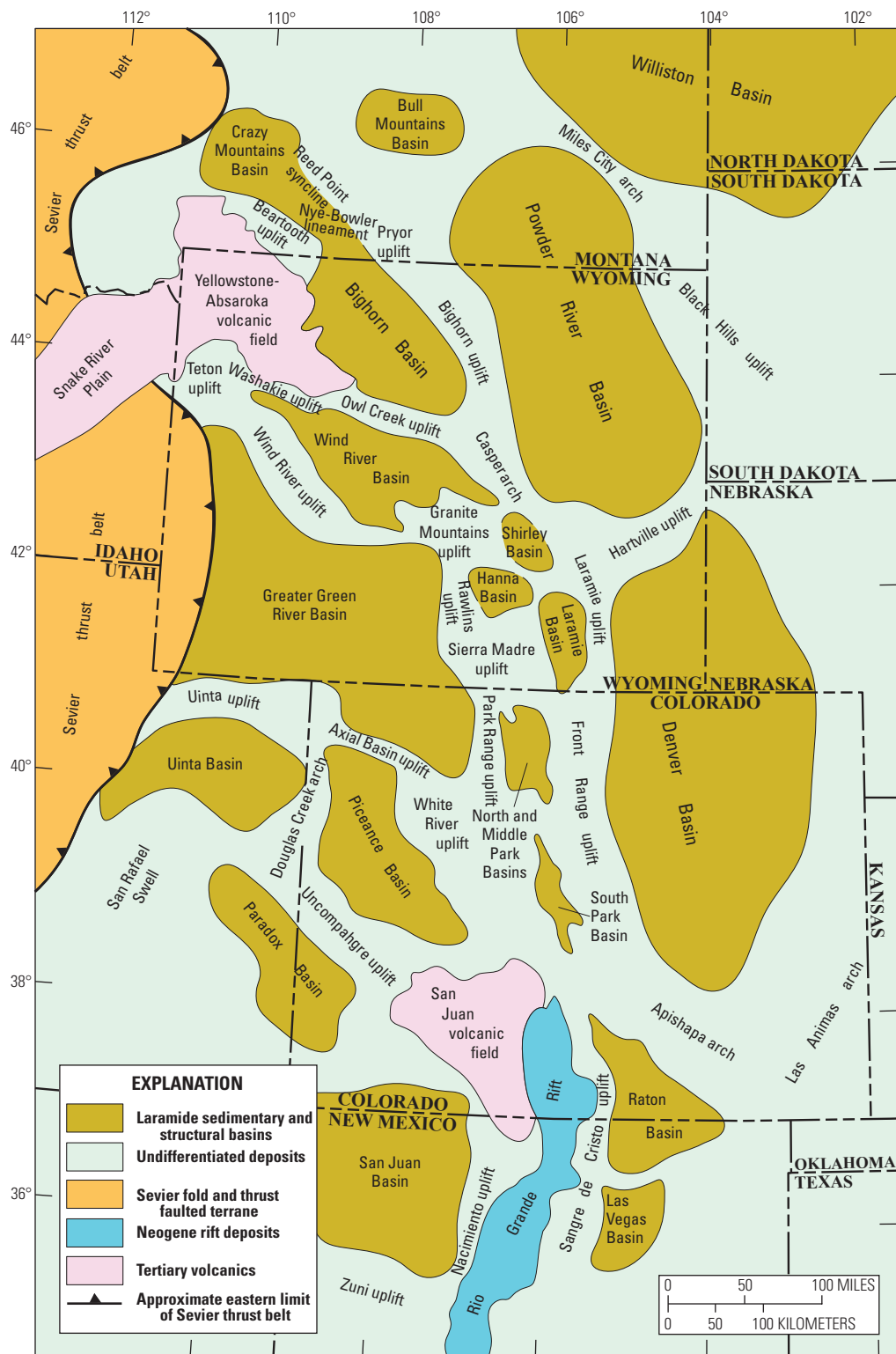


Figure 1. Map of the Rocky Mountain region showing locations of Laramide basins.

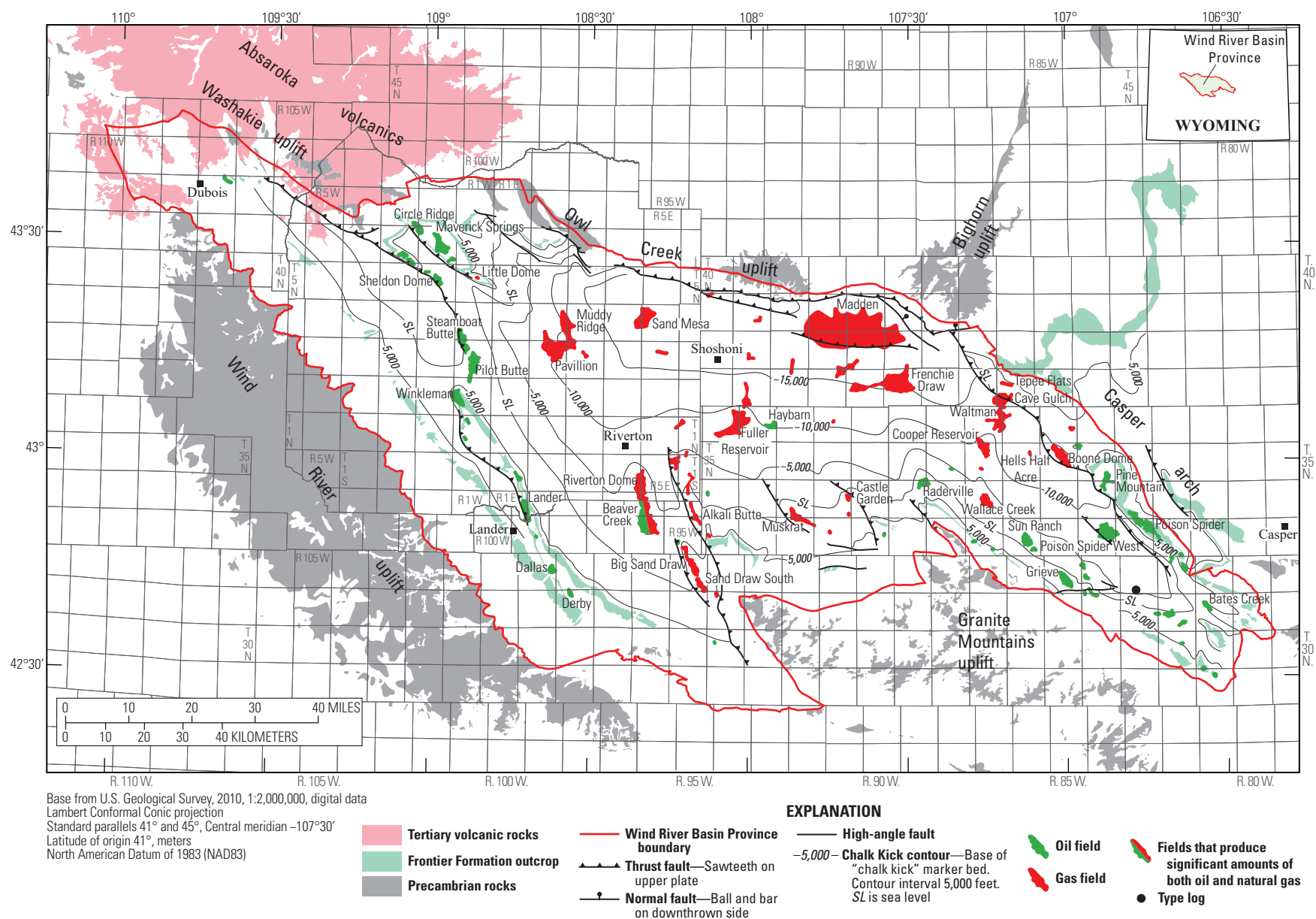
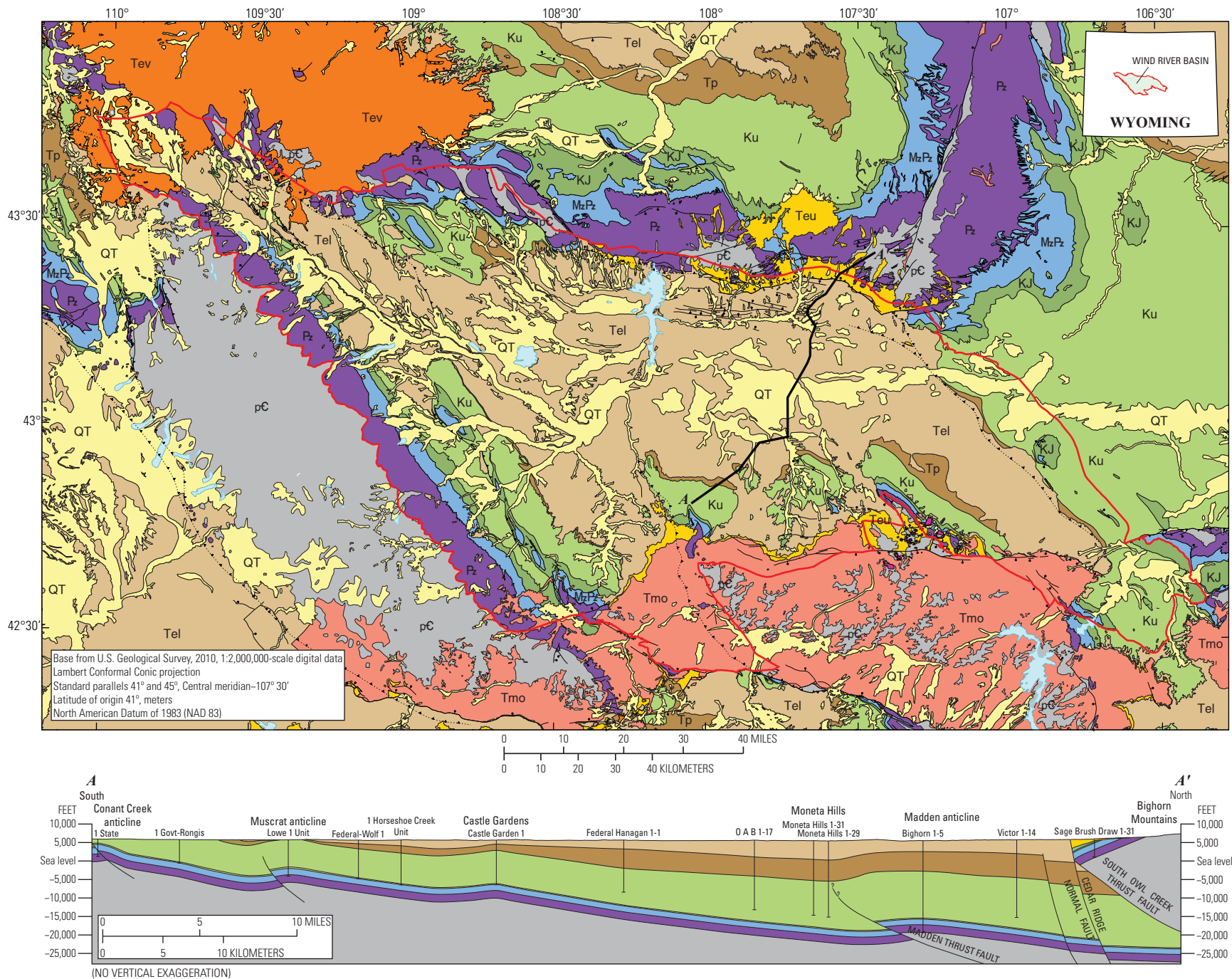


Figure 2. Index map of the Wind River Basin in central Wyoming showing major structural elements, and principal oil and gas producing fields. Structure contours drawn at base of the chalk kick marker bed of the Niobrara interval of the Cody Shale. Contour interval is 5,000 feet. SL, sea level.



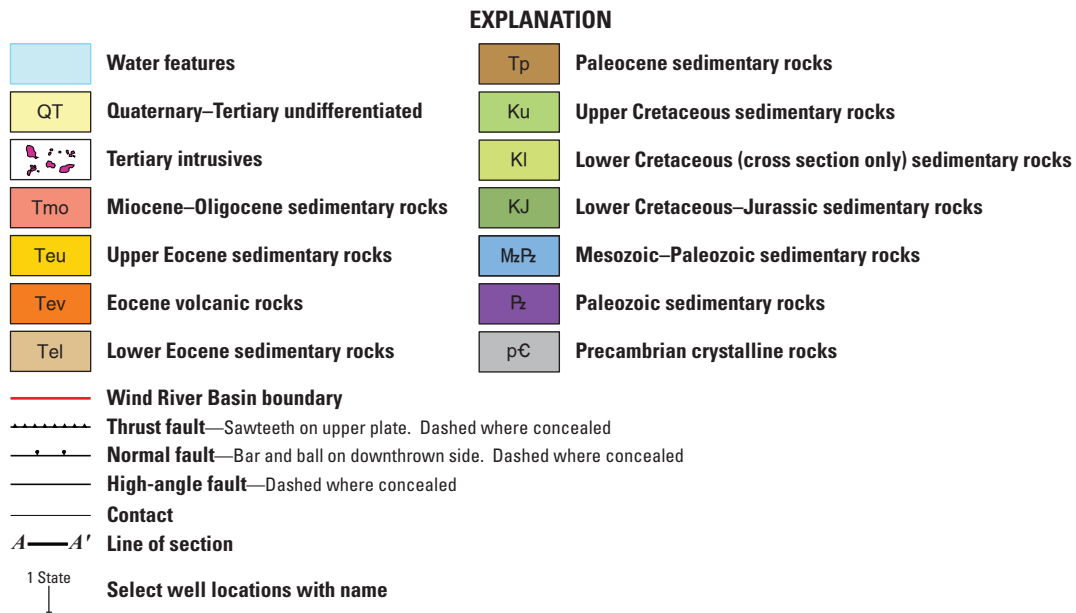
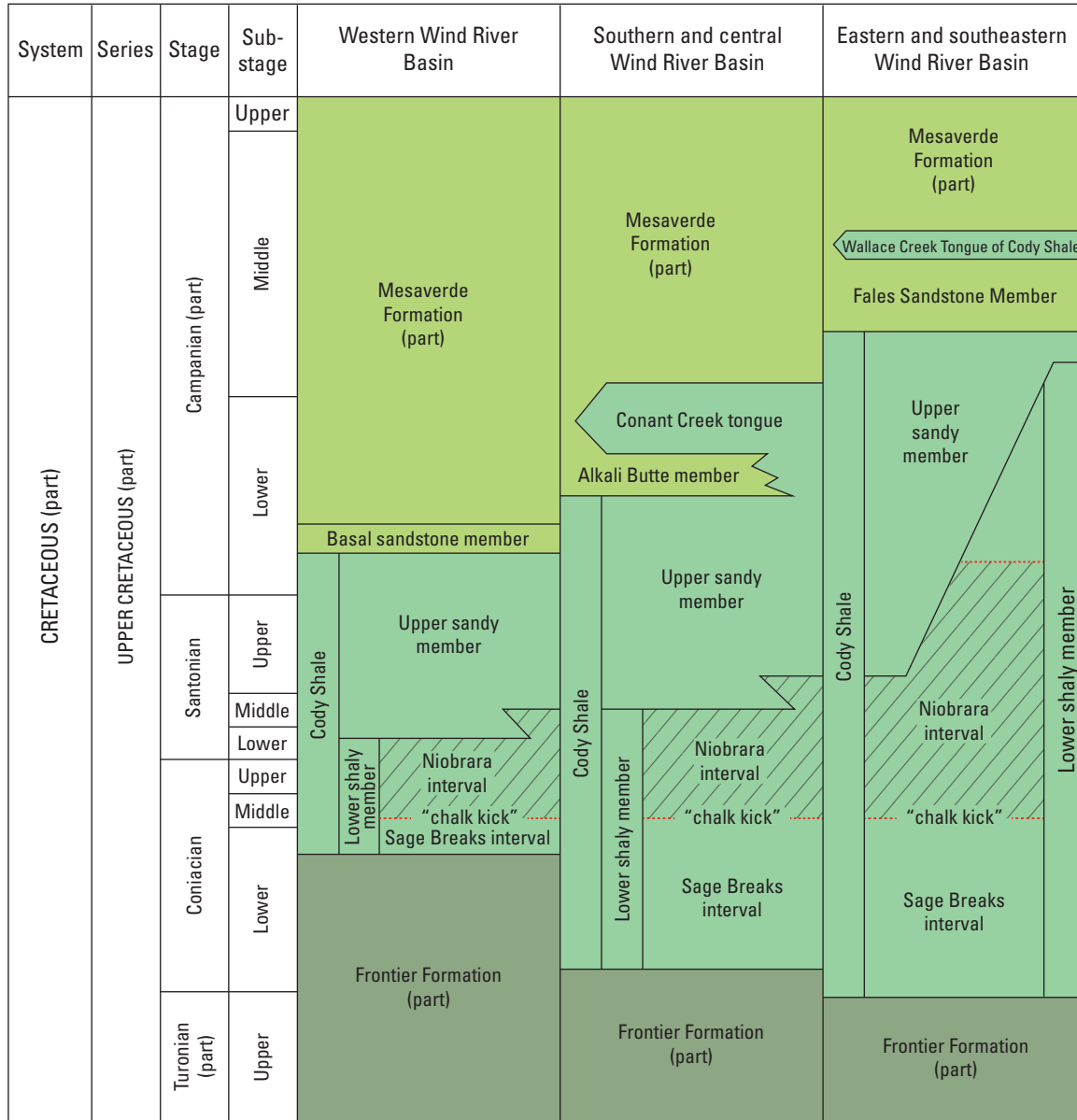


Figure 3. (above and previous page) Generalized geologic map and cross section of the Wind River Basin in central Wyoming. Geologic map modified from Love and Christiansen (1985), and Green and Drouillard (1994). Cross section segment north of the Madden anticline is modified from Ray and Keefer (1985).

and Sutherland (2009). Additional sources of geologic data include structural cross sections and (or) subsurface maps by Biggs and Espach (1960), Keefer (1970), Skeen and Ray (1983), Sprague (1983), Ray and Keefer (1985), Dunleavy and Gilbertson (1986), Cardinal and others (1989), Blackstone (1990), Natali and others (2000), and Montgomery and others (2001). The structure contour map is at a scale of 1:500,000, the contour interval is 1,000 ft, and the datum is mean sea-level (sheet 1). The overburden map was constructed by using the well data (Finn, 2019) and by calculating the difference between the surface elevation and the structure contours. This map is at a scale of 1:500,000, and the contour interval is 1,000 ft (sheet 2).



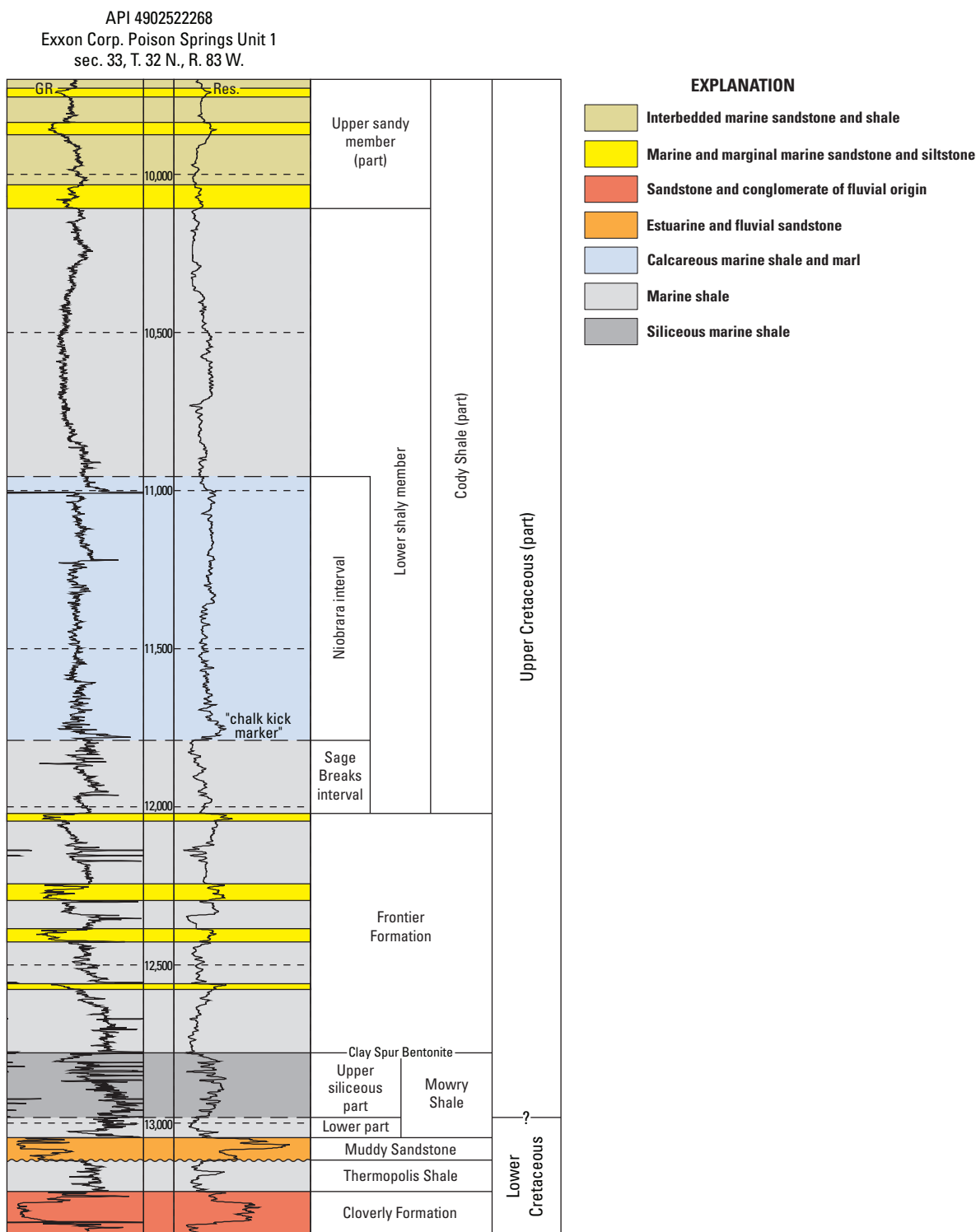


Figure 5. Type log of Lower and lowermost Upper Cretaceous rocks in the southeastern part of the Wind River Basin. GR, gamma ray log; Res., resistivity log. Location shown in figure 2. From Finn (2017).

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