

Structure Contour and Overburden Maps of the Niobrara Interval of the Upper Cretaceous Cody Shale in the Bighorn Basin, Wyoming and Montana

By Thomas M. Finn

Pamphlet to accompany

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Conversion Factors

U.S. customary units to International System of Units

Multiply	By	To obtain
Length		
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
Area		
square mile (mi ²)	2.590	square kilometer (km ²)

Datum

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above sea level.

Structure Contour and Overburden Maps of the Niobrara Interval of the Upper Cretaceous Cody Shale in the Bighorn Basin, Wyoming and Montana

By Thomas M. Finn

Introduction

The Bighorn Basin is a complex intermontane sedimentary and structural basin that formed during the Laramide orogeny (Late Cretaceous through early Eocene). The basin is nearly 180 miles (mi) long, 100 mi wide, and covers approximately 10,400 square miles (mi²) in north-central Wyoming and a small part of south-central Montana (fig. 1). The basin is bounded on the northeast by the Pryor uplift, on the east by the Bighorn uplift, and on the south by the Owl Creek uplift (fig. 2). The northern boundary of the Bighorn Basin is formed by the Nye-Bowler lineament, a regional anticlinal trend extending about 60 mi east-southeast from the northern part of the Beartooth uplift to the Pryor uplift (figs. 1 and 2). This trend consists of a series of highly faulted anticlines and domes that Wilson (1936) interpreted to overlie a left-lateral basement shear zone. The Nye-Bowler lineament forms the structural divide separating the Bighorn Basin to the south from the Reed Point syncline and Crazy Mountains Basin to the northwest (fig. 1). The northwest and west margins are formed by the Beartooth uplift and Absaroka volcanic field, respectively (fig. 2). The basin margins are characterized by highly folded and faulted sedimentary rocks that range from Paleozoic to Paleocene in age; whereas, the central part of the basin is covered by nearly flat-lying lower Eocene and undifferentiated Tertiary and Quaternary rocks that mask the structure in the central part of the basin (fig. 3).

According to Hewett and Lupton (1917) and Berry (1952), oil was first discovered in the Bighorn Basin in 1884 emanating from the Upper Cretaceous Mowry Shale at a seep near what is now the Bonanza field along the east margin of the basin (fig. 2). This discovery led to the drilling of several shallow wells in the area; however, no oil or gas was found in commercial amounts (Berry, 1952; Biggs and Espach, 1960). It wasn't until 1906 that the first commercial hydrocarbon production in the Bighorn Basin was established from Cretaceous reservoirs at Garland field followed by the discovery of Greybull field in 1907 (Biggs and Espach, 1960; Fox and Dolton, 1996) (fig. 2). Since then, many important conventional oil and gas resources have been discovered and produced from reservoirs ranging in age from Cambrian to

Tertiary (fig. 2) (De Bruin, 1993; Fox and Dolton, 1996). In addition, a potential continuous (unconventional) basin-centered gas accumulation may be present in Cretaceous reservoirs in the deeper parts of the basin (Ryder, 1987; Spencer, 1987; Surdam and others, 1997; Johnson and Finn, 1998a; Johnson and others, 1999; Finn and others, 2010). Numerous authors (see for example, Burtner and Warner, 1984; Hagen and Surdam, 1984; Meissner and others, 1984; Ryder, 1987; and Fox and Dolton, 1989, 1996) have suggested that various Upper Cretaceous marine shales are the principal hydrocarbon source rocks for many of these accumulations. With recent advances in horizontal drilling and multi-stage hydraulic fracture stimulation, equivalent Upper Cretaceous marine source rock intervals, in particular the Niobrara Formation, have become important shale gas or shale oil objectives in other Rocky Mountain Laramide basins (Matthews, 2011; Sonnenberg, 2011; Williams and Lyle, 2011; Durham, 2012a,b, 2013; Taylor and Sonnenberg, 2014; Hawkins, 2016). In the Bighorn Basin, the Niobrara is represented by shales, calcareous shales, marls, siltstones, and sandstones in the lower shaly member of the Cody Shale (Finn, 2014, 2019) (fig. 4).

The maps presented in this report were constructed as part of a project carried out by the U.S. Geological Survey (USGS) to better characterize the geologic framework of potential undiscovered continuous (unconventional) oil and gas resources of the Niobrara interval of the Upper Cretaceous Cody Shale in the Bighorn Basin in north-central Wyoming and south-central Montana (sheets 1 and 2). The structure contour map is drawn at the base of the "chalk kick" marker bed, a distinctive zone or peak on resistivity logs in the lower 50–300 feet (ft) of the Cody Shale that according to Finn (2014, 2019) represents the base of the Niobrara interval in the Bighorn Basin (fig. 5). This horizon was selected because it is easily identified on most well logs, is present throughout the basin, and was identified by the author in about 1,250 wells (Finn, 2020). In addition, about 20 wells in the deeper parts of the basin that penetrated the upper part of the Niobrara interval or overlying Cody Shale were utilized by projecting thicknesses from nearby wells and estimating a pick for the "chalk kick" marker (Finn, 2020). Along the shallow margins of the basin, many wells (about 60) were used that were drilled through the Niobrara interval; however, in these wells, the

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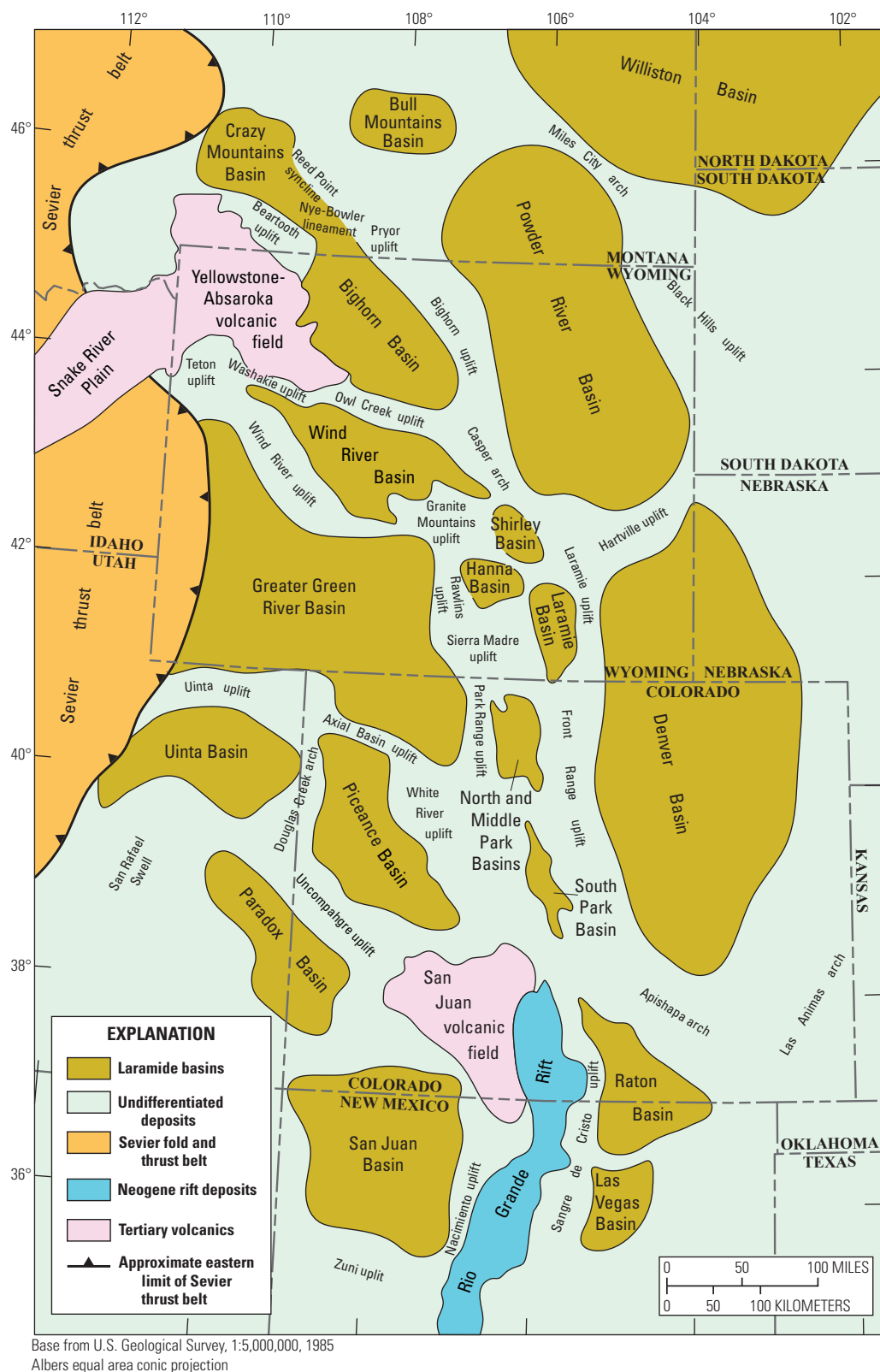
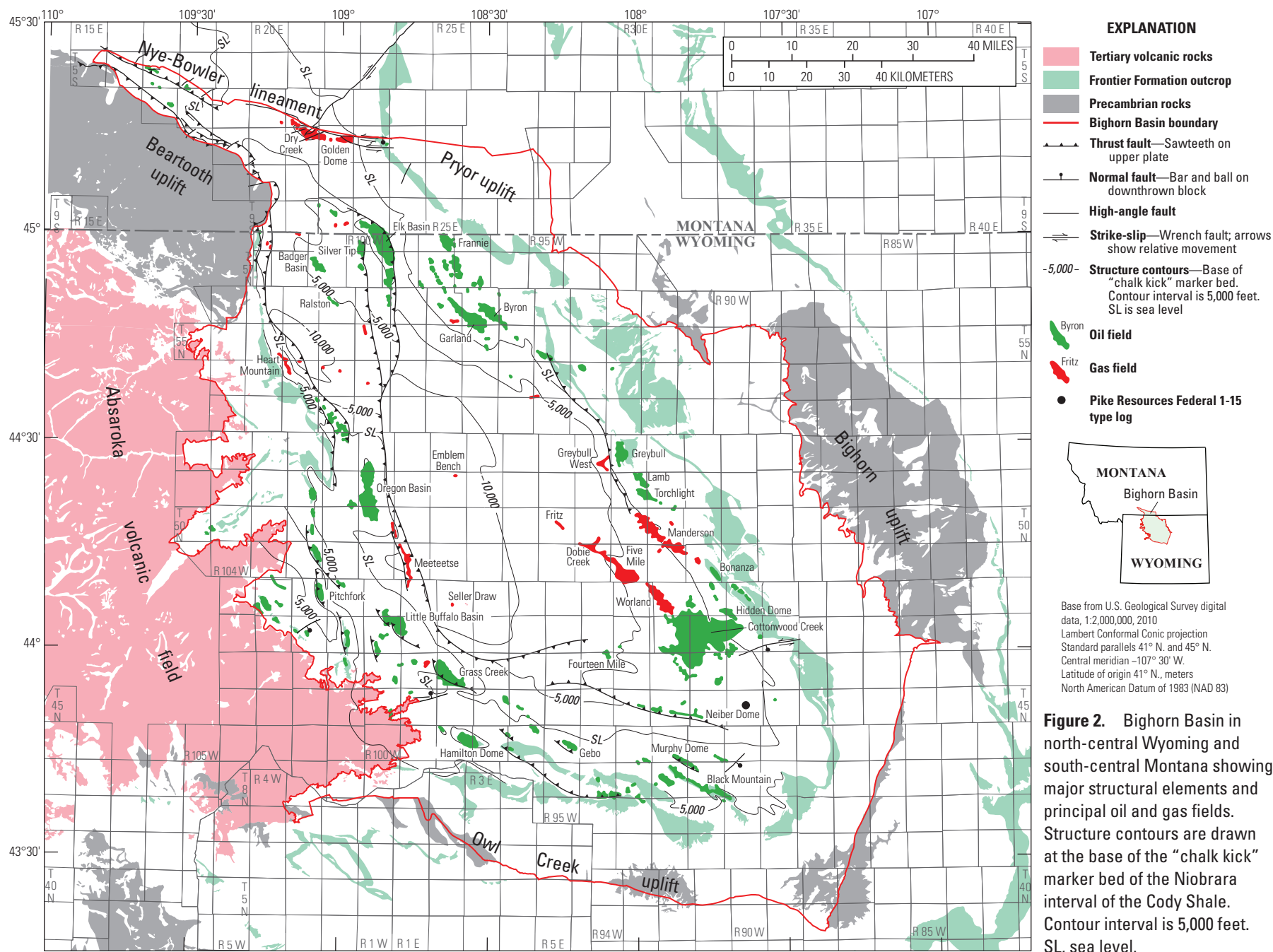


Figure 1. Rocky Mountain region showing locations of Laramide basins.



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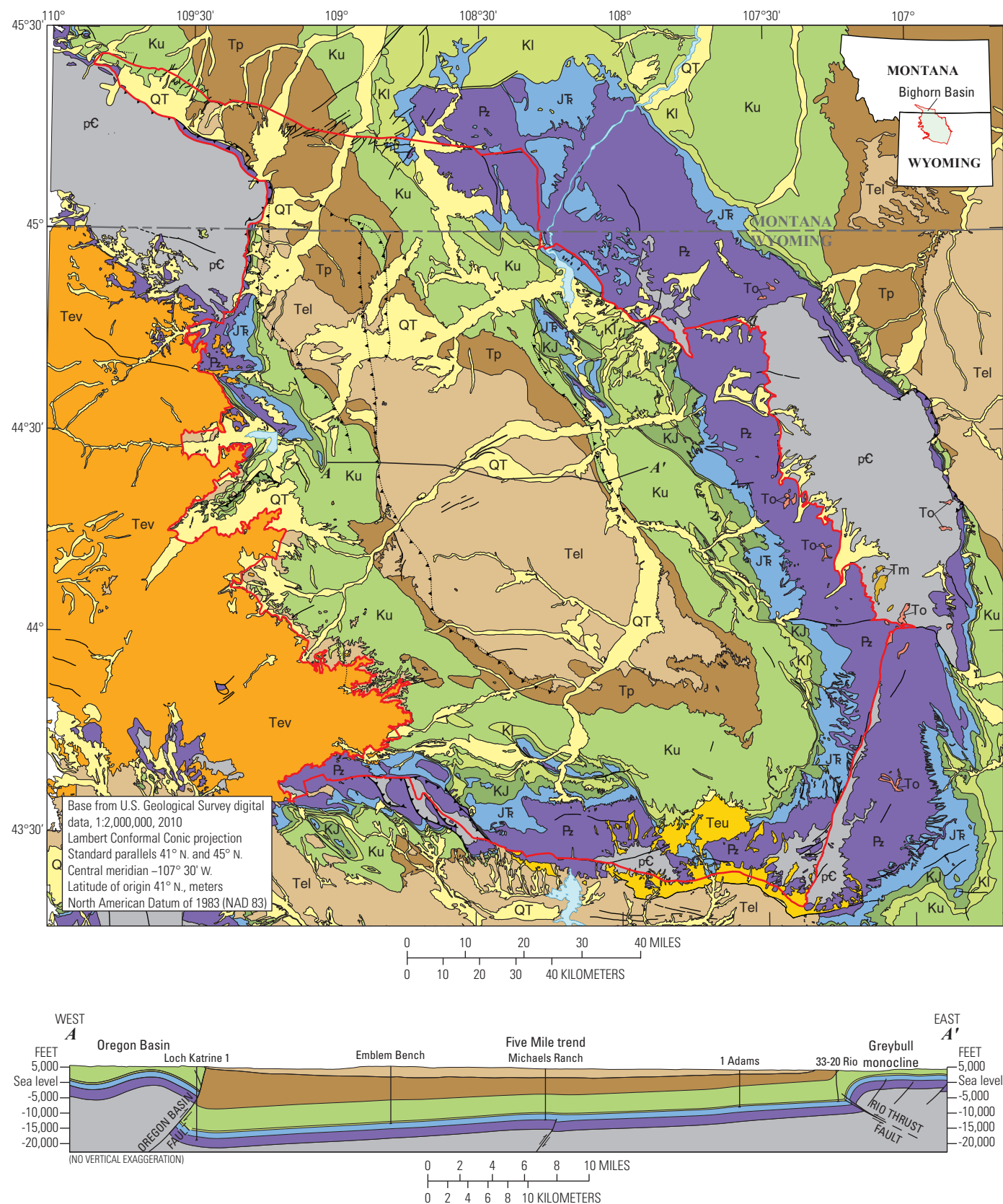


Figure 3 (above and following page). Bighorn Basin in north-central Wyoming and south-central Montana. Geologic map modified from Love and Christiansen (1985), Green and Drouillard (1994), and Raines and Johnson (1995).

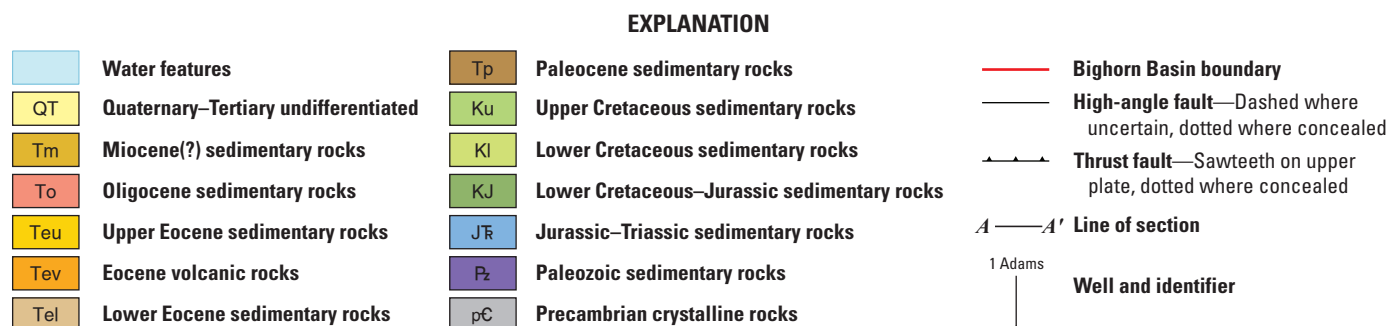


Figure 3.—Continued Bighorn Basin in north-central Wyoming and south-central Montana. Geologic map modified from Love and Christiansen (1985), Green and Drouillard (1994), and Raines and Johnson (1995).

System	Series	Stage	Sub-stage	Northern Bighorn Basin, Montana	Northern Bighorn Basin, Wyoming	Western and southwestern Bighorn Basin	Eastern and southeastern Bighorn Basin			
Cretaceous (part)	Upper Cretaceous (part)	Campanian (part)	Middle (part)	Judith River Formation (part)	Mesaverde Formation (part)	Mesaverde Formation (part)	Mesaverde Formation (part)			
				Parkman Sandstone	Claggett Member		Claggett Member			
				Claggett Shale						
			Lower	Eagle Sandstone	Lower member		Lower member			
				Telegraph Creek Formation	Upper sandy member		Upper sandy member	Upper sandy member		
		Santonian	Cody Shale	Cody Shale		Cody Shale			Cody Shale	Upper sandy member
					Middle					
		Coniacian	Lower	Lower	Cody Shale	Cody Shale	Cody Shale	Cody Shale	Upper sandy member	
				Upper						
				Middle						
			Lower	Lower shaly member	Niobrara interval "chalk kick"	Niobrara interval "chalk kick"	Niobrara interval "chalk kick"	Niobrara interval "chalk kick"	Lower shaly member	Sage Breaks interval
				Sage Breaks interval	Sage Breaks interval	Sage Breaks interval	Sage Breaks interval			
Turonian (part)	Upper	Frontier Formation (part)	Frontier Formation (part)	Frontier Formation (part)	Frontier Formation (part)	Frontier Formation (part)				

Figure 4. Upper Cretaceous Cody Shale and associated strata in the Bighorn Basin, Wyoming and Montana.

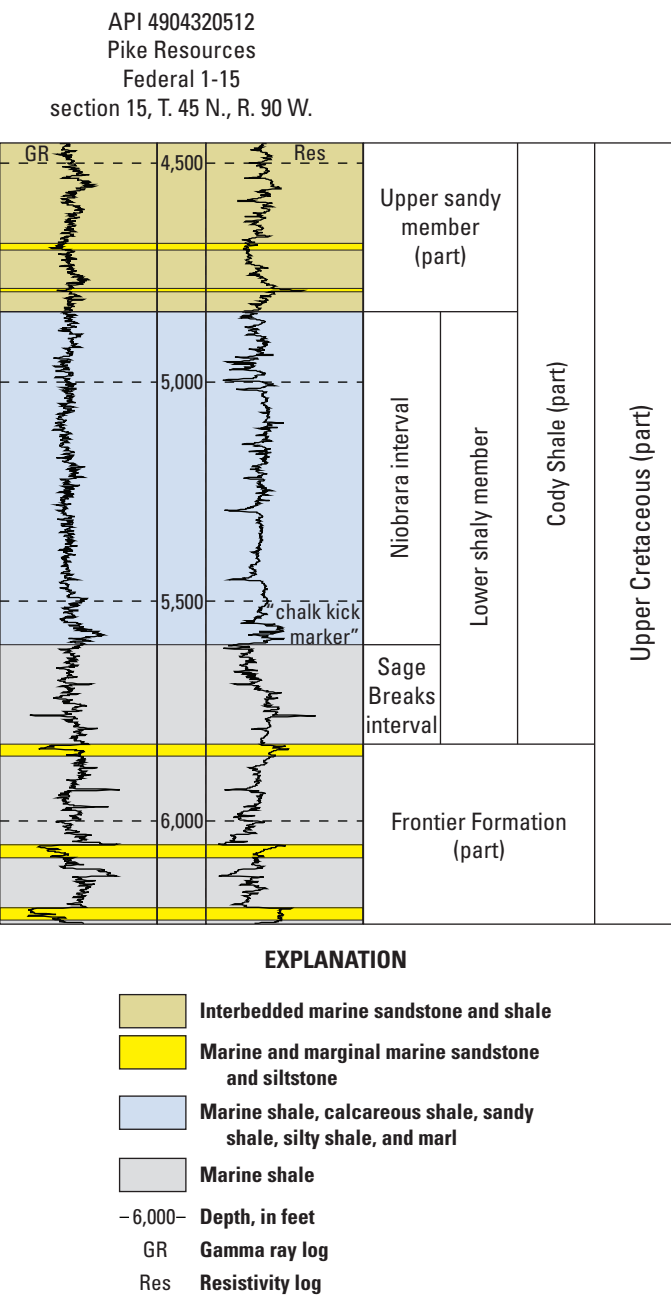


Figure 5. Niobrara interval and associated strata in the Pike Resources Federal 1-15 well in the southeastern part of the Bighorn Basin. Location shown in figure 2. GR, gamma ray log; Res., resistivity log.

logging run was started in the underlying Frontier Formation. In these wells, the depth to the “chalk kick” marker was extrapolated up hole to estimate its depth to aid in contouring (Finn, 2020). The structure contouring is based on interpretation of the well data (Finn, 2020), surface geologic mapping by Love and Christiansen (1985), Green and Drouillard (1994), Raines and Johnson (1995), and Vuke and others (2007). Additional sources of geologic data include structural cross sections and (or) subsurface mapping by Stewart (1975), Stone (1975, 1983a,b, 1985a,b, 2004a,b), Tonnsen (1985), Ver Ploeg (1985), Bikis and Anderson (1986), Blackstone (1986a,b), Grauman and others (1986), Cardinal and others (1989), Johnson and Finn (1998b). The structure contour map is at a scale of 1:500,000, the contour interval is 1,000 ft, and the datum is mean sea level (sheet 1). The overburden map was constructed by using the well data (Finn, 2020) and calculating the difference between the surface elevation and the structure contours. This map is at a scale of 1:500,000, and the contour interval is 1,000 ft (sheet 2).

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