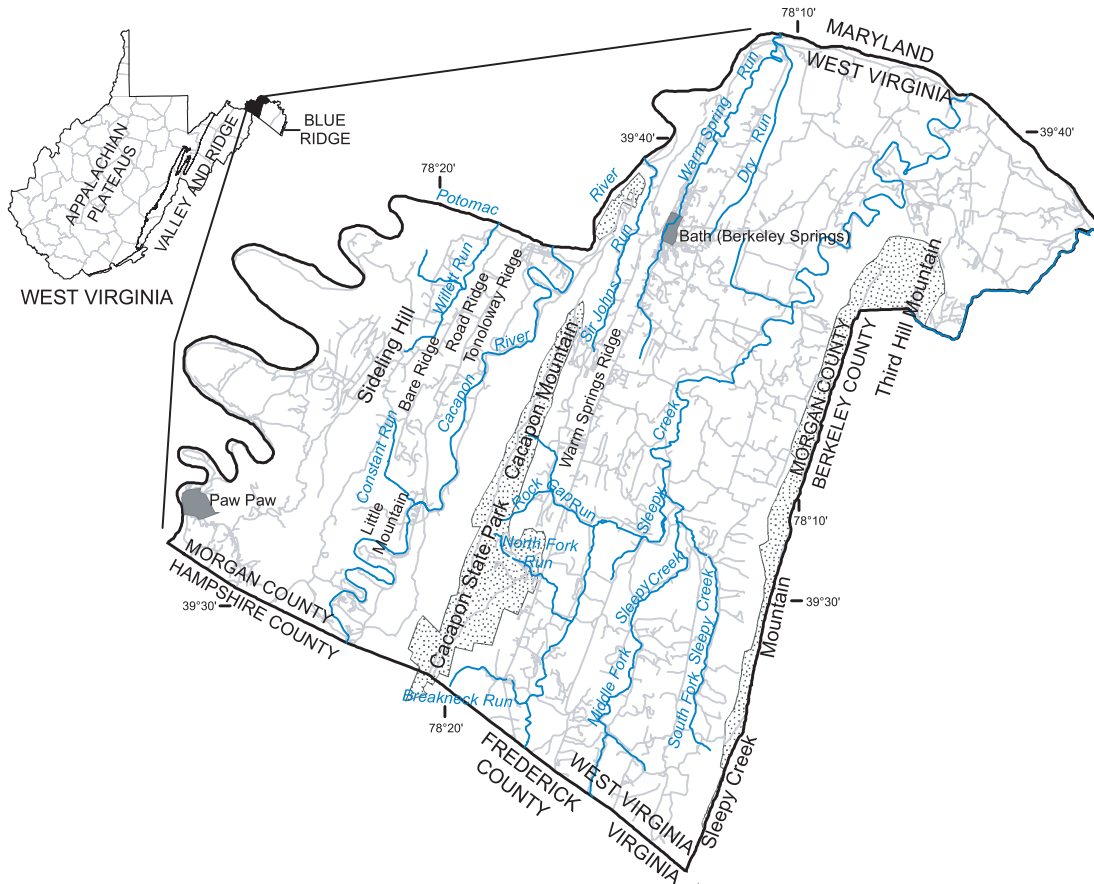


Prepared in cooperation with the West Virginia Conservation Agency and the Eastern Panhandle Conservation District

Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia



Scientific Investigations Report 2006-5198

Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia

By Carol J. Boughton and Kurt J. McCoy

Prepared in cooperation with the West Virginia Conservation Agency and the Eastern Panhandle Conservation District

Scientific Investigations Report 2006–5198

U.S. Department of the Interior
U.S. Geological Survey

U.S. Department of the Interior
DIRK KEMPTHORNE, Secretary

U.S. Geological Survey
P. Patrick Leahy, Acting Director

U.S. Geological Survey, Reston, Virginia: 2006

For product and ordering information:
World Wide Web: <http://www.usgs.gov/pubprod>
Telephone: 1-888-ASK-USGS

For more information on the USGS—the Federal source for science about the Earth, its natural and living resources, natural hazards, and the environment:
World Wide Web: <http://www.usgs.gov>
Telephone: 1-888-ASK-USGS

Any use of trade, product, or firm names is for descriptive purposes only and does not imply endorsement by the U.S. Government.

Although this report is in the public domain, permission must be secured from the individual copyright owners to reproduce any copyrighted materials contained within this report.

Suggested citation:
Boughton, C.J., and McCoy, K.J., 2006, Hydrogeology, aquifer geochemistry, and ground-water quality in Morgan County, West Virginia: U.S. Geological Survey Scientific Investigations Report 2006–5198, 56 p.

Contents

Abstract.....	1
Introduction.....	1
Purpose and Scope	1
Description of Study Area	1
Methods of Investigation.....	3
Previous Studies	5
Hydrogeology.....	6
Geology.....	6
Hydrology	9
Relation of Geology to Ground-Water Flow.....	9
Aquifer Properties	10
Water Levels.....	15
Aquifer Geochemistry and Ground-Water Quality	17
Geochemical Properties.....	17
Water-Quality Characteristics	17
pH.....	17
Specific Conductance.....	17
Temperature.....	20
Alkalinity	21
Water Hardness	22
Dissolved Oxygen.....	22
Turbidity.....	22
Inorganic Constituents.....	22
Major Ions	23
Bromide and Fluoride	23
Iron and Manganese	27
Trace Elements.....	29
Nutrients.....	29
Nitrate and Nitrite	29
Ammonia Species	30
Phosphorous and Orthophosphate	30
Radon-222.....	31
Bacteria	32
Summary and Conclusions.....	33
Acknowledgments	34
References Cited.....	34
Appendix 1. Well-construction and aquifer properties for selected wells, Morgan County, West Virginia	38
Appendix 2a-d. Water-quality data values for wells, Morgan County, West Virginia.....	41
Appendix 3a-c. Water-quality data values for springs, Morgan County, West Virginia.....	54

Figures

1.-3. Maps showing:	
1. Location and physical features of study area	2
2. Location of wells and springs sampled in Morgan County, West Virginia	4
3. Geology of Morgan County, West Virginia, including major structural features and lines of section	7
4. Generalized geologic section of the Great Cacapon and Paw Paw quadrangles, Morgan County, West Virginia	8
5. Generalized geologic section of Stotlers Crossroads quadrangle, Morgan County, West Virginia	8
6. Conceptual model of ground-water flow through a faulted and fractured anticlinal ridge bordered by a synclinal ridge	11
7. Block diagram showing possible circulation of flow to a warm spring located near the crest of an anticline at the intersection of two faults.....	11
8. Map showing locations and magnitudes of transmissivities, Morgan County, West Virginia	14
9. Map showing ground-water-level altitudes and depth to static water levels in Morgan County, West Virginia	16
10.-15. Boxplots showing:	
10. Distribution of pH values in ground-water samples from hydrogeologic units located in Morgan County, West Virginia	18
11. Relation between concentration of pH and alkalinity in ground-water samples from Morgan County, West Virginia.....	18
12. Relation between concentration of pH and dissolved silica in ground-water samples from Morgan County, West Virginia	19
13. Distribution of specific conductance in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.....	19
14. Distribution of temperature in ground-water samples from springs and wells located in Morgan County, West Virginia	20
15. Distribution of temperature in ground-water samples from hydrogeologic units located in Morgan County, West Virginia	21
16. Diagrams showing relative geochemical composition of ground water from springs and wells, by hydrogeologic unit, Morgan County, West Virginia.....	24
17.-22. Boxplots showing:	
17. Distribution of concentrations of iron in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.....	27
18. Distribution of concentrations of manganese in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.....	28
19. Relation between concentrations of iron and manganese in ground-water samples from Morgan County, West Virginia.....	28
20. Relation between concentrations of zinc and aluminum in ground-water samples from Morgan County, West Virginia.....	30
21. Distribution of concentration of radon-222 in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.....	31
22. Relation between the ratio of casing length to well depth and the occurrence of bacteria in wells located in Morgan County, West Virginia.....	33

Table

1. Comparison of historical transmissivity and specific-capacity data (Kozar and Mathes, 2001) with results from this study for geologic units occurring in Morgan County, West Virginia12

Conversion Factors, Datum, and Abbreviations

Multiply	By	To obtain
Length		
inch (in.)	25.4	millimeter (mm)
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
Area		
square mile (mi ²)	2.590	square kilometer (km ²)
Flow rate		
gallon per minute (gal/min)	0.06309	liter per second (L/s)
inch per year (in/yr)	25.4	millimeter per year (mm/yr)
Radioactivity		
picocurie per liter (pCi/L)	0.037	becquerel per liter (Bq/L)
Specific capacity		
gallon per minute per foot [(gal/min)/ft]	0.2070	liter per second per meter [(L/s)/m]
Transmissivity		
foot squared per day (ft ² /d)	0.09290	meter squared per day (m ² /d)

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as follows:

$$^{\circ}\text{F}=(1.8\times^{\circ}\text{C})+32$$

Vertical coordinate information is referenced to the National Geodetic Vertical Datum of 1929 (NGVD 29).

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above the vertical datum.

Transmissivity: The standard unit for transmissivity is cubic foot per day per square foot times foot of aquifer thickness [(ft³/d)/ft²ft]. In this report, the mathematically reduced form, foot squared per day (ft²/d), is used for convenience.

Specific conductance is given in microsiemens per centimeter at 25 degrees Celsius (μS/cm at 25°C).

Concentrations of chemical constituents in water are given either in milligrams per liter (mg/L) or micrograms per liter (μg/L).

This page intentionally blank.

Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia

By Carol J. Boughton and Kurt J. McCoy

Abstract

Private and public wells throughout Morgan County, W. Va., were tested to determine aquifer hydraulic, geochemical, and water-quality characteristics. The entire study area is located in the Valley and Ridge Physiographic Province, a region of complex geologic structure and lithology. Aquifers in the study area are characterized by thin to thick bedded formations with interbedding among the various limestones, shales, sandstones, and siltstones that are folded into a series of steeply dipping north-south trending anticlines and synclines. Zones of ground-water production typically consist of one to two fracture sets, with little to no production from unfractured bedrock matrix. Measurements of transmissivity range from 2 to 1,490 feet squared per day, with the larger transmissivities occurring near bedding contacts and in zones with cross-faulting or jointing. Ground water flows from recharge areas in the uplands to local drainages and to deeper flow systems that appear to be controlled by regional geologic structure. The overall flow direction is from south to north within the study area.

Ground water within the study area is predominantly a calcium-bicarbonate type water reflecting contact with carbonate rocks. Sodium-bicarbonate and calcium-magnesium-sulfate end-members also exist, with many samples exhibiting mixing, which may be the result of flow between the differing rock types or within units containing both carbonate rocks and shales. Values of water-quality characteristics that were greater than U.S. Environmental Protection Agency drinking-water standards included radon-222, pH, turbidity, iron, manganese, aluminum, and total- and fecal-coliform and *Escherichia coli* (*E. coli*) bacteria. Concentrations of radon-222 were detected in all samples from all units, with the largest concentrations (1,330 and 2,170 picocuries per liter) from the Clinton Formation.

Introduction

Economic expansion and population growth from the Baltimore, Md., and Washington, D.C., metropolitan corridor to Morgan County has led to increased construction of homes and businesses with commensurate demands on available

ground-water supplies. Additional sources of good-quality water are needed to allow for continued population and economic growth and to provide alternative supplies in the event of a drought such as that during 2001–02. The U.S. Geological Survey (USGS), in cooperation with the West Virginia Conservation Agency and the Eastern Panhandle Conservation District, conducted a study to assess ground-water conditions to provide technical data that could be used as a basis for management of Morgan County water resources and future planning. Additionally, USGS assessed water-quality and geochemical characteristics of springs in cooperation with West Virginia University.

Purpose and Scope

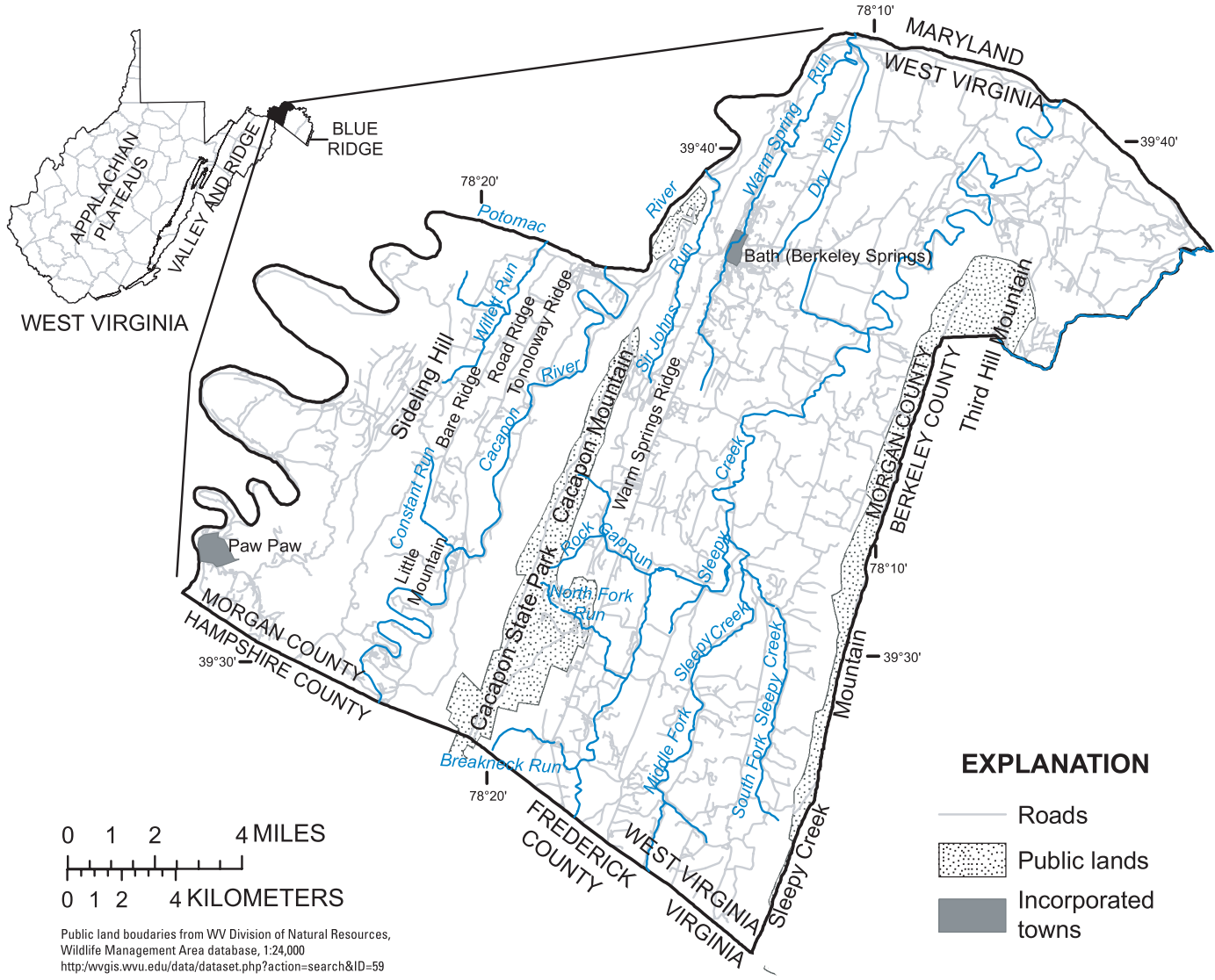
This report describes hydrogeologic, geochemical, and water-quality characteristics of aquifers and geochemical and water-quality characteristics of select springs located in Morgan County, W. Va. The aquifer characteristics include the hydrogeologic, physical, and geochemical setting that may affect the observed quality. Geochemical, water-quality, and hydraulic data were collected over a 1-year period beginning in the spring of 2004 through the spring of 2005.

Description of Study Area

Morgan County occupies approximately 231 mi² in the eastern panhandle of West Virginia (fig. 1). The county was formed in 1820 from parts of Hampshire and Berkeley Counties and was named in honor of General Daniel Morgan, a prominent soldier of the Revolution (Holmes, 1995). The Potomac River forms the county's western and northern boundaries. Hampshire County, W. Va., and Frederick County, Va., form the southwestern and southeastern boundaries, respectively. Berkeley County, W. Va., bounds Morgan County to the east.

Morgan County lies in the Valley and Ridge Physiographic Province as defined by Fenneman (1946) and is characterized by a series of north-northeast trending anticlines and synclines, which form prominent valleys and ridges. The major physiographic features of much of Morgan County are steep-sloped mountains alternating with narrow valleys.

2 Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia



Public land boundaries from WV Division of Natural Resources, Wildlife Management Area database, 1:24,000
<http://wvgis.wvu.edu/data/dataset.php?action=search&ID=59>

Roads from USGS DLG data files, 1:24,000
<http://wvgis.wvu.edu/data/dataset.php?action=search&ID=100>

Hydrography from Natural Resource Analysis Center (NRAC) at WVU NHD mapping, 1:24,000
<http://wvgis.wvu.edu/data/dataset.php?action=search&ID=235>

County and town boundaries from USGS DLG data from 1:24,000 topographic maps
<http://wvgis.wvu.edu/data/dataset.php?action=search&ID=77>

Figure 1. Location and physical features of study area.

The exception is the eastern part of the county, which is characterized by rolling terrain transitioning to the Great Valley along the Route 9 corridor.

The climate of Morgan County is affected both by prevailing continental storm tracks from the west and maritime effects from the Atlantic Ocean (Doll and others, 1963). Precipitation in Morgan County falls as winter snow and as spring, summer, and fall rains. Weakened hurricanes can produce considerable rain during the late summer and early fall as the storm tracks flow northeastward along the Allegheny front. The mean annual precipitation as measured at Cacapon State Park is approximately 39 in/yr (National Oceanic and Atmospheric Administration, 2002) and ranges from approximately 36 in/yr in the valleys to 44 in/yr along the Sleepy Creek Mountain divide (Hobba and others, 1972).

In addition to the Potomac River, the Cacapon River and Sleepy Creek are two relatively large perennial streams within the county. Both streams flow in a northerly direction and drain into the Potomac River. The Sleepy Creek watershed is entirely within the county.

Smaller streams include Warm Springs Run, Breakneck Run, and Dry Run. The Potomac River bounds Morgan County on the west and north. All other streams tend to flow in a northerly direction parallel to major geologic structural controls and discharge into the Potomac River. Tributaries tend to follow trellis geomorphic patterns within structurally controlled basins. Sir Johns Run and Rock Gap Run are among exceptions that flow through interbasin water gaps controlled by lineaments (Lessing and others, 1991).

Historically, land use in Morgan County has been that of the rural landscape. Grimsley (1916) reported that Morgan County was one of the important peach and apple orchard sections of the state, exporting large quantities of those fruits. Additionally, the county had up to 25 canneries, mostly for tomatoes, another significant agricultural product (Grimsley, 1916). Two incorporated towns existed in 1910, Berkeley Springs (Bath) and Paw Paw, having populations of 864 and 725, respectively (Grimsley, 1916). The leading industry was sand and glass, derived from the relatively pure, silica-rich Oriskany Sandstone that occurs in the county (Ice, 1933).

As reported in 1980, approximately 81 percent of land use in Morgan County was forest, 17 percent was agricultural land, and 1 percent was urban or built-up land (McColloch and Lessing, 1980). However, landscape and demographics within the county have changed appreciably during the past 20 years. In 1980, the Morgan County population was 10,606 (U.S. Department of Commerce, 1991). According to the 2000 census, the population had increased to 14,943, an approximate 40 percent increase in 20 years. During 1990–2000, the population increased by 23.2 percent

(U.S. Department of Commerce, 2005). Former agricultural and forested land is being urbanized to accommodate the increasing population.

Methods of Investigation

Well sampling sites were selected on the basis of (1) spatial distribution throughout the county and within representative geologic formations; (2) well age, in that wells that were recently completed were preferred, whenever possible; and (3) landowner's willingness to provide permission for access to existing wells.

Well-completion reports from the West Virginia Bureau for Public Health for 565 wells completed in Morgan County during 1997–2002 were reviewed for potential sampling-site suitability. The well-completion reports were evaluated for completeness and lithologic accuracy based upon knowledge of the geology within the county. Recently completed wells were preferred to ensure that the wells were properly grouted and sealed, thereby minimizing the potential for surface contamination along an open well bore (West Virginia Bureau for Public Health, 1984). After initial screening, less than one-half of the reports were considered to represent potentially appropriate well sites. Wells were excluded on the basis of incorrect lithologic descriptions and inadequate detail.

Potential wells were evaluated to provide spatial distribution throughout the county and among geologic formations. Most of the wells that were sampled were domestic wells used daily by the homeowners. Others were used seasonally or only for watering yards or for watering livestock, and some were public-supply wells.

Hydrogeologic units were sampled in an east-west cross section through the county with most aquifers having representative sites to the northern and the southern limits of the county to maximize areal distribution of sampling sites. Researchers at West Virginia University selected the spring sites for sampling. The USGS sampled these in cooperation with the university researchers who did a detailed assessment of springs in the study area. The resulting ground-water sampling network of 91 wells and 8 springs provides information about the variability of hydrogeologic characteristics, ground-water quality, and aquifer geochemistry within and among the water-bearing units (fig. 2).

Ground water was sampled using the techniques described in Wilde and others (1999) by purging each well with its existing pump. Springs were sampled with a pump as close to the point of emergence as possible. Measurements of pH, specific conductance (SC), dissolved oxygen (DO), water temperature, and turbidity in water from the pumping well or spring were recorded until stability criteria were met.

4 Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia

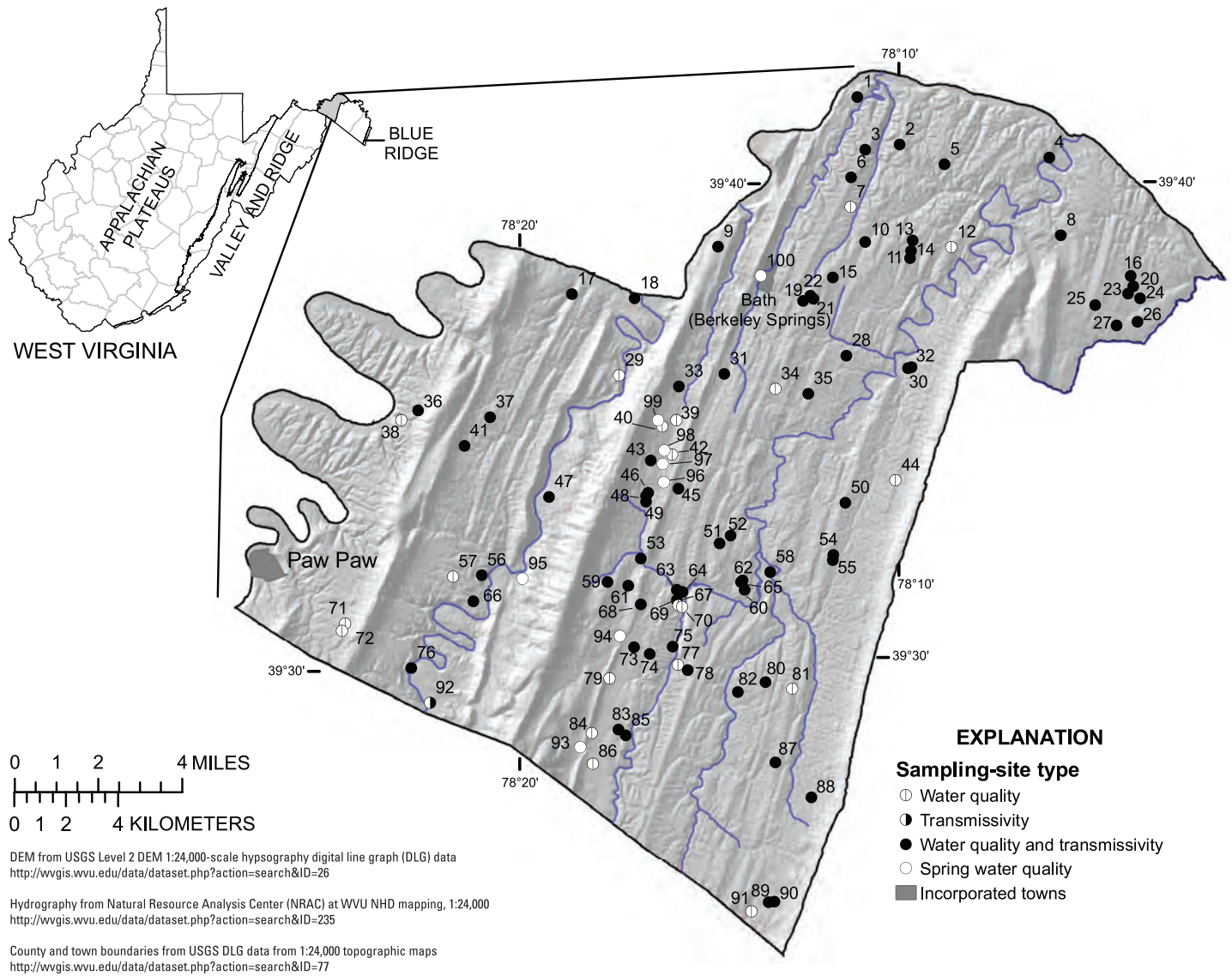


Figure 2. Location of wells and springs sampled in Morgan County, West Virginia.

Then samples were collected and preserved in the field. Samples were shipped to the USGS National Water Quality Laboratory (NWQL) in Lakewood, Colo., where they were analyzed for concentrations of calcium, magnesium, sodium, potassium, chloride, sulfate, fluoride, iron, manganese, silica, total dissolved solids, hardness, and nutrients (nitrogen and phosphorous species) (American Public Health Association, 1998; Fishman, 1993; Fishman and Friedman, 1989; Patton and Truitt, 1992). Subsets of ground-water samples were analyzed for the trace elements bromide, barium, arsenic, aluminum, zinc, strontium, cadmium, lead, antimony, and nickel (Faires, 1993; Fishman and Friedman, 1989; Garbarino, 1999; Garbarino and Struzeski, 1998) and for the naturally occurring radioactive gas radon-222 (American National Standards Institute, 2004).

Carbonate alkalinity, bicarbonate alkalinity, and total alkalinity were measured in the field using the techniques described in Wilde and Radtke (1998). Total-coliform, fecal-coliform, and *Escherichia coli* (*E. coli*) bacteria were plated on site, incubated, and counted using techniques described in Myers and Wilde (2003).

Previous Studies

A detailed, systematic assessment of ground-water conditions throughout Morgan County has not been published previously. Grimsley (1916) provided an account of the historical and industrial development, the physiography, and the geology of the eastern panhandle counties, including Morgan County. The warm springs now known as Berkeley Springs in the town of Bath, W. Va., were prized by native Americans before the arrival of white settlers and George Washington noted them in his journal as early as 1747. According to Grimsley's account (1916), the springs were known for their uniform water temperature from 21 to 23°C throughout the year. Erskine (1948) documented seasonal temperatures and discharge from Berkeley Spring, Ziler Spring, and Breakneck Run Springs in his discussion of the principal springs in West Virginia.

Hoffman (1962) anticipated that the water needs of individual states, including West Virginia, might parallel the future national trend. He further suggested that West Virginia needed a close appraisal of its water resources and water problems due, in part, to its proximity to the eastern seaboard. According to Doll and others (1963), inadequate hydrologic data for planning purposes and increasing competition for water were among the water problems facing West Virginia.

Morgan County has been included in several regional studies. Doll and others (1963) described the water resources of the Potomac River Basin in West Virginia. Three wells in Morgan County, one from the Oriskany Sandstone and two from the Chemung Group, were identified as yielding 50 gal/min or more. They noted that artesian conditions occur in the deep-lying sandstone beds, reducing the amount of lift required but not necessarily producing large well yields. The

largest reported concentration of total iron from shale aquifers was 10.0 mg/L and from limestone aquifers was 5.0 mg/L. Water from shale aquifers tended to be more acidic than that from limestone aquifers. Within the Potomac River Basin, water from shallow, karstic limestone aquifers tended to have larger concentrations of nitrate and chloride at some locations. The one sample from Morgan County had much smaller concentrations of nitrate and chloride than other sites discussed (Doll and others, 1963).

Hobba and others (1972) published a compendium of water-resources information for the Potomac River Basin. Sandstone and carbonate-rock aquifers within the basin were characterized as having the largest supplies available, owing to fractures and solution cavities. Throughout the Potomac River Basin, carbonate aquifer yields are highly variable, from 5 to over 500 gal/min, and depend upon the well depth and local geologic structure. The Oriskany Sandstone yields were variable depending upon whether the unit is well jointed, cemented, or friable. Low water tables may occur beneath ridges underlain by permeable zones of the Oriskany Sandstone as ground-water drains from the unit. In Morgan County, only one well, near Paw Paw, W. Va., was identified as producing from alluvium. Ground-water quality throughout the Potomac River Basin was found to be variable because of the complexity of the lithology and structure (Hobba and others, 1972).

Schwietering (1981) expanded on the work of Hobba and others (1972), describing ground-water conditions and aquifers on a statewide scale for West Virginia. Geohydrologic units were described in detail; however, few data were published specific to Morgan County except for a further discussion of Berkeley Springs.

Hobba (1985) describes ground-water quantity and quality in Hardy, Hampshire, and western Morgan Counties. These limited published data illustrate wide ranges in hydrogeologic and water-quality characteristics for sites in Morgan County.

Ferrell (1987) and Mathes and others (1998) provide summaries of available ground-water quality data, and Kozar and Mathes (2001) provide a compendium of aquifer characteristics throughout West Virginia. However, few data were available for Morgan County in these studies.

Cooper and Hanna (1988) compiled instantaneous discharge and temperature data for surface-water stations throughout West Virginia for 1945–75. Cacapon River near Great Cacapon, W. Va., was the only site located in Morgan County for which data were available.

Hobba and others (1979) discussed the geology, hydrology, and geochemistry of the Berkeley Springs area in their study of thermal springs in the Appalachians. McColloch (1986) produced a compendium of springs of West Virginia that summarizes available data for 14 springs in Morgan County, including Hancock Station Spring, Michael Farm Spring, Berkeley Springs, Robert Milborn Spring, Fahren Spring, Mitchum Farm Spring, Coolfont Spring, Whisner Farm Spring, Neely Farm Spring, Ziler Spring, Cacapon State

Park Spring, Ridge Hatchery Spring, O.B. Overmale Springs, and an unnamed spring.

Hydrogeology

Ground water is stored in intergranular spaces, in bedrock fracture features, and in solutionally enhanced fractures within Morgan County. The intergranular space within the bedrock matrix is referred to as primary porosity. Ground water does not flow readily through the intergranular spaces in bedrock because of a lack of connectivity and of cementation of voids. Ground water flows more readily through bedrock with secondary porosity. Secondary porosity occurs in fracture features, including faults, joints, and partings at contact zones and in solutionally enhanced features, which have much larger effective porosity. The structural and depositional geology controls the directional orientation and size of the fractures.

Geology

The rocks of Morgan County (fig. 3) are of Ordovician (Oswego and Juniata Formations), Silurian (Tuscarora Sandstone, Clinton Group, McKenzie Formation, Williamsport Sandstone, Wills Creek Formation, Tonoloway Formation), Devonian (Helderberg Group, Oriskany Sandstone, Onesquethaw Group, Marcellus Formation, Mahantango Formation, Harrell Shale, Brallier Formation, Chemung Group, and Hampshire Formation), and Mississippian (Pocono Group), ages (Cardwell and others, 1968). The rocks were folded and faulted extensively during the formation of the Appalachian Mountains, resulting in a series of northeast-southwest trending anticlines and synclines with beds of varying width (figs. 4 and 5). The beds plunge slightly to the north. Underlying resistant rocks such as sandstones form the ridges and less resistant rocks such as shales typically form the valleys (Hobba and others, 1972).

Lithologic descriptions in well-completion reports and other available mapping are approximate on a site-specific basis; therefore, throughout this report, undifferentiated references to the Brallier Formation will include the Brallier Formation and the Harrell Shale and references to the Tonoloway Formation will include the Tonoloway Formation, Wills Creek Formation, and Williamsport Sandstone. The Needmore Shale and Huntersville Chert are mapped as the Onesquethaw Group (Cardwell and others, 1968), but the section is so narrow in the study area that it is not mappable at the 1:250,000 scale.

Deposits of unconsolidated alluvial material typically occur near major streams and consist primarily of clay and silt, with some discontinuous lenses of sand and gravel. Alluvial deposits occasionally occur on terraces, but are typically

unsaturated. Significant saturated thicknesses are generally found only near major streams (Hobba and others, 1972). No wells were identified as producing from the alluvium in this study; therefore, the alluvial aquifer system is not discussed further in this report.

The Pocono Group is gray, hard, massive sandstone with some shale beds. The thickness of the Pocono Group is as much as 1,030 ft (Cardwell and others, 1968). It is underlain by the Hampshire Formation, a maroon to reddish-brown siltstone and shale in the upper 300 to 400 ft of the formation and fine-grained, well-cemented, hard, massive red sandstone with some shale in the lower part. The thickness of the Hampshire Formation is as much as 3,350 ft (Cardwell and others, 1968). The Chemung Group is typically a rusty-yellowish brown. The upper formation is thin-bedded, hard, blocky sandstone underlain principally by thin-bedded siltstone and shale. The thickness of the Chemung Group is as much as 2,115 ft (Cardwell and others, 1968). A massive sandstone member occurs at its base in the easternmost areas (Hobba and others, 1972). The Pocono Group, Hampshire Formation, and Chemung Group occur extensively in both the western and eastern parts of the county.

The Brallier Formation is a greenish-gray shale with siltstone and sandstone that constitutes approximately 35 percent of the formation. The thickness of the Brallier Formation is as much as 1,300 ft (Cardwell and others, 1968). The Harrell Shale is black shale that occurs in the lower 50 to 400 ft of this formation. The thickness of the Harrell Shale is zero to 50 ft (Cardwell and others, 1968). Lying below the Harrell Shale, the Mahantango Formation consists of brown sandy shales with occasional sandstone beds and concretions found near the contact with the underlying Marcellus Formation. The thickness of the Mahantango Formation is as much as 1,200 ft (Cardwell and others, 1968). The Marcellus Formation is black carbonaceous, clayey, thinly laminated, fissile shale containing pyrite and minor amounts of siltstone (Reger, 1924). The thickness of the Marcellus Formation is as much as 200 ft (Cardwell and others, 1968). Thinly bedded and 20 to 40 ft thick, the Purcell Member of Pennsylvanian limestone (Cardwell and others, 1968) occurs approximately 50 to 150 ft from the base (Reger, 1924), and large limestone concretions are known to occur near the base of the Marcellus Formation (Hobba and others, 1972).

The Onesquethaw Group, consisting of the Needmore Shale and Huntersville Chert, occurs below the Marcellus Formation. The Needmore Shale is a predominately dark gray or green, calcitic, mostly nonfissile, gamma-ray-emitting shale. It includes the Black Beaver Dam Shale member and grades westward into the Huntersville Chert. The Tioga bentonite occurs near the top of the Needmore Shale. The thickness of the Needmore Shale ranges from 75 ft to 140 ft (Cardwell and others, 1968).

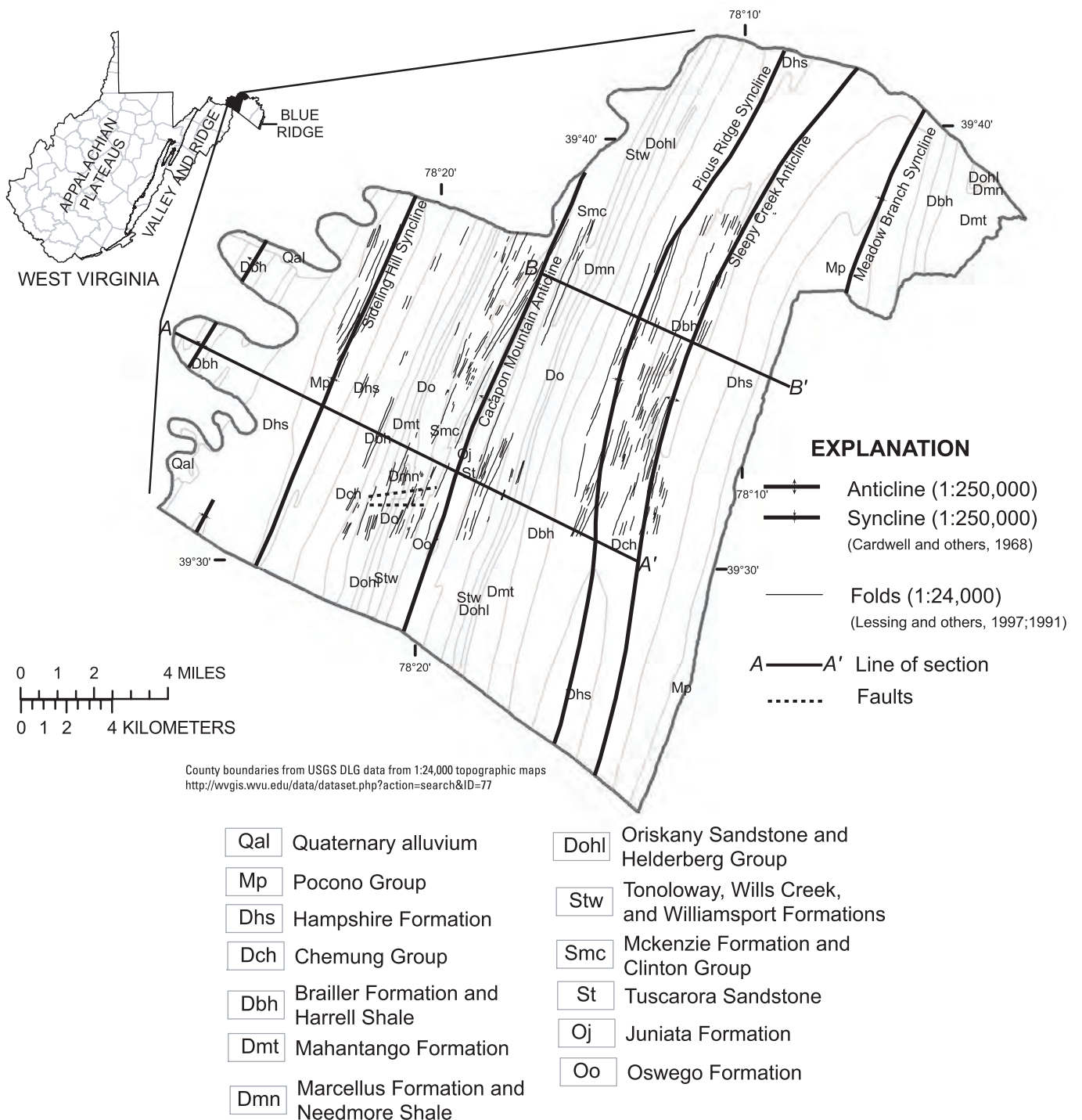


Figure 3. Geology of Morgan County, West Virginia, including major structural features and lines of section. Geology modified from Lessing and others (1997), Lessing and others (1991), and Cardwell (1968).

8 Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia

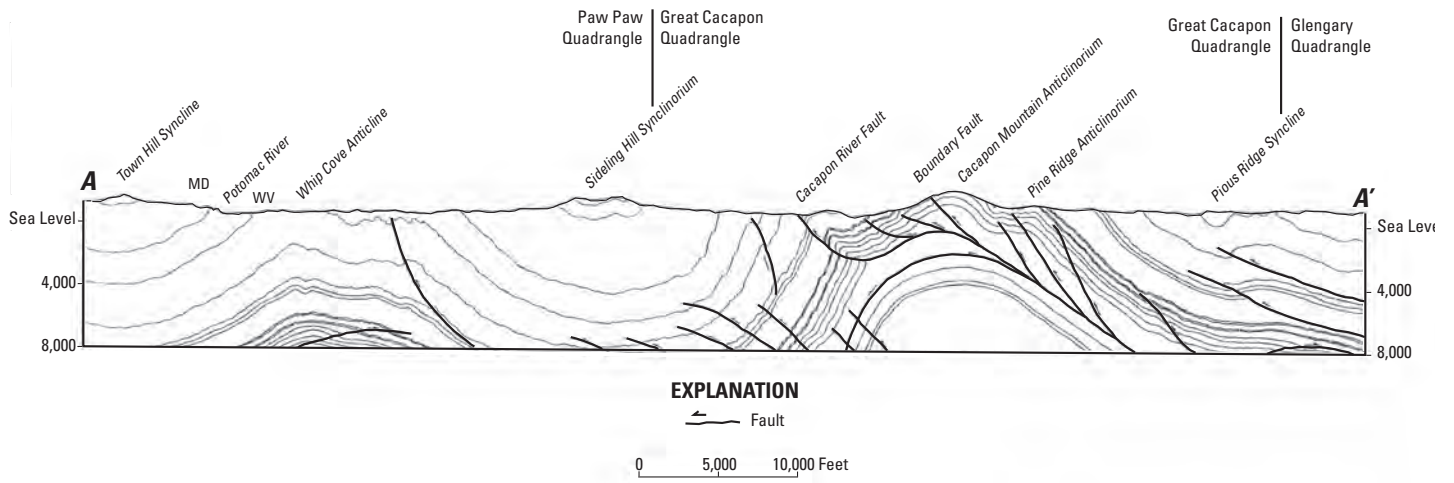


Figure 4. Generalized geologic section of the Great Cacapon and Paw Paw quadrangles, Morgan County, West Virginia. Modified from Lessing and others (1997). (Line section of A-A' as shown on figure 3.)

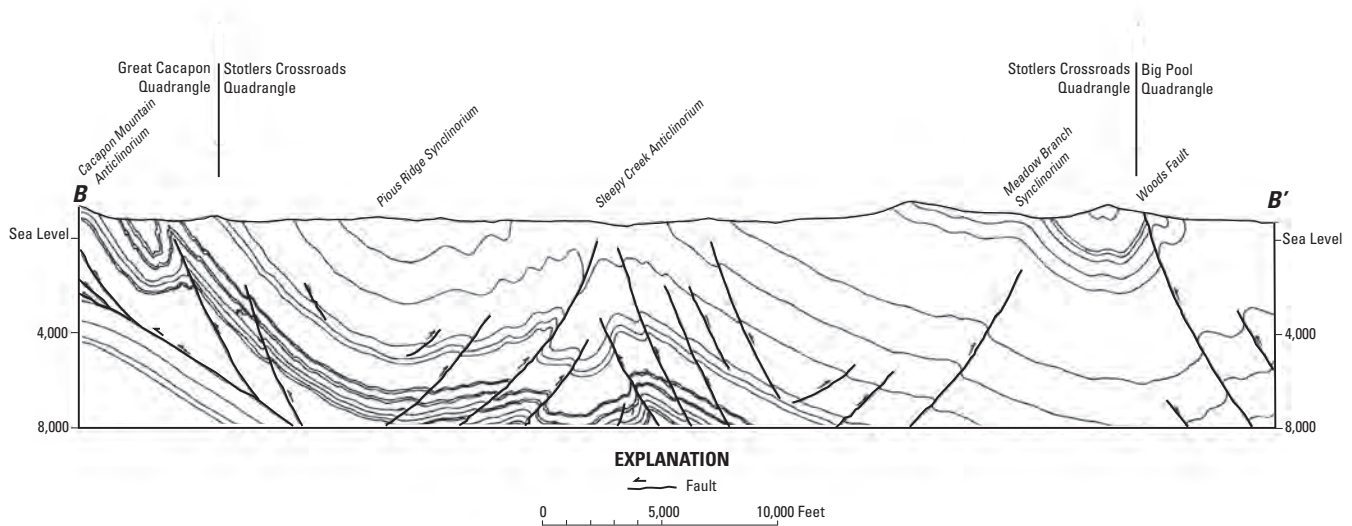


Figure 5. Generalized geologic section of Stotlers Crossroads quadrangle, Morgan County, West Virginia. Modified from Kulander and others (1995). (Line of section B-B' as shown on figure 3.)

The other member of the Onesquethaw Group is the Huntersville Chert, which grades from a nearly pure, slightly calcitic or dolomitic chert (interbedded with the Needmore Shale), to a limestone westward. That limestone, commonly referred to as the Onondaga, includes the glauconitic $[(K,Na)(Fe^{3+},Al,Mg)_2(Si,Al)_4O_{10}(OH)_2]$ Bobs Ridge Sandstone member (Cardwell and others, 1968). Although the Onesquethaw (Onondaga) Group is narrow and may not constitute a potentially significant aquifer, its mineralogy may affect ground-water quality where it does occur. The gamma-ray emissions may be associated with radon in drinking water. Glauconite is used as a water softener, so its occurrence could account for zones of naturally occurring soft water.

The Oriskany Sandstone is a medium- to thick-bedded brown to white sandstone lying between two layers of soft calcareous chert. The thickness of the Oriskany Sandstone ranges from 12 ft to 275 ft (Cardwell and others, 1968). Cementation of these rocks varies from very hard silica cement to softer carbonate cement to friable and devoid of cement depending upon the location. The upper chert layer ranges in thickness from 0 to 100 ft. In Morgan County, the lower Shriver Chert is very thin or absent (Hobba and others, 1972).

The Helderberg Group varies from thin-bedded, light-colored, cherty limestone in the upper part to thin-bedded shale in the middle part to massive-bedded relatively pure limestone in the lower part. The thickness of the Helderberg Group is as much as 170 ft (Cardwell and others, 1968). The underlying Tonoloway Formation consists of gray, flaggy, thin beds containing some silica and dolomite and grading downward into flaggy limestone and calcareous shales of the Wills Creek Shale. Red or green sandy shales and one hard sandstone bed occur at the base (Hobba and others, 1972). The thickness of the Tonoloway Formation is as much as 590 ft and the thickness of the Wills Creek Shale is as much as 400 ft (Cardwell and others, 1968).

The McKenzie Formation is a bluish-gray, wavy-bedded, flaggy limestone and limy shale, whereas the underlying Clinton Group consists of red shale with thin beds of sandstone and occasional limestone layers. The thickness of the McKenzie Formation is as much as 240 ft (Cardwell and others, 1968). The Tuscarora Sandstone is thick-bedded, extremely hard, resistant, and is colored gray to pink. By way of contrast, the "red beds" of the Juniata Formation are thin-bedded block sandstone and shale. In places, this formation is underlain by thick-bedded gray Oswego Sandstone (Hobba and others, 1972). The thickness of the Tuscarora Sandstone is as much as 100 ft; of the Juniata Formation is as much as 320 ft; and of the Oswego Sandstone is as much as 300 ft (Cardwell and others, 1968).

Hydrology

The study area is predominantly underlain by folded, faulted, and fractured sedimentary shale, carbonate, and sandstone rocks that form the aquifer systems. Recharge occurs

throughout the study area, particularly in the upland areas. Discharge occurs locally to streams and springs, then regionally to the Potomac River. Ground-water discharge, including spring discharge, provides base flow to the streams.

Relation of Geology to Ground-Water Flow

The flow of ground water in the study area is complicated because of the structural and lithologic complexity of the Valley and Ridge Physiographic Province. The bedrock generally has small primary permeability in the rock matrix and much larger secondary permeability in fractures, along bedrock contacts, and in solutionally enhanced zones of preferential flow. Similar to the hydrogeology of the Appalachian Plateaus Physiographic Province (Wyrick and Borchers, 1981), wells that do not penetrate fractures or solution channels will not produce usable amounts of water. Flow paths in Morgan County are complicated by the complex system of steeply dipping, north-south trending anticlines and synclines that form the topography typical of the Valley and Ridge Physiographic Province.

Well yields in primarily limestone rock along the axes of anticlines are found to be appreciably larger than those along axes of synclines (Clark and others, 1976). Siddiqui and Parizek (1972) found that variation in structural control is not a statistically significant factor in the productivity of wells not located on fracture traces. Ground-water flow in Morgan County may occur through fracture sets that parallel anticlinal axes. Where deep vertical faults or horizontal thrust faults intersect these fractures, they provide conduits for potentially large well yield.

McCoy and others (2005a, 2005b) found in Berkeley and Jefferson Counties that discrete fracture zones within the bedrock yield water to the majority of wells tested. Over 500 recent (1997–2004) well-completion reports for Morgan County were examined during the course of this study. Not all completion reports clearly indicated the water-production zones encountered. The completion reports that indicate location of production zones show a clear pattern of production from one to two discrete zones, sometimes described as being at contacts between units and sometimes as apparent fracture zones within a given bedrock unit. There was no indication of usable, continuous production along the well bore from either fractures or bedrock matrix. Although many fractures intersect a borehole, only a few may yield usable quantities of ground water. Aquifer yields are dependent upon water produced from fractures, faults, or along contact zones.

Ground water flows from areas with higher hydraulic head toward areas of lower hydraulic head. The preferential flow path followed will be that which has the largest transmissivity or ability to transmit water. Toth (1963) describes an idealized model for flow systems in most field areas with negligible local topographic relief; however, in areas with pronounced local relief, such as the Valley and Ridge Province, it is possible that only local systems develop (Freeze and Cherry, 1979).

Within the study area, the ground water flows from recharge areas in the uplifted geologic features such as Cacapon Mountain and Sleepy Creek Mountain toward the valleys and then discharges to the north toward the Potomac River. Locally, ground water flows to springs, streams, or nearby pumping wells that intercept the flow and cause minor, localized perturbations in the flow path.

Transmissivity, typically, is the largest downdip along bedding planes or fractures parallel to bedding planes, although intersecting faults or fracture zones may act as drains and route flow across structural features. The folded and uplifted bedrock of the Appalachian Mountains may serve as a barrier to flow between major tributary basins such as the Cacapon and Warm Springs Run watersheds except where flow occurs along bedding planes or through other fractured and faulted features (fig. 6).

Cross-faulting occurs near the town of Bath, where the Berkeley Springs arise (Hobba and others, 1979). Ground water contributing to the Berkeley Springs is thought to be derived from recharge to and deep flow through the Oriskany Sandstone. Water budget, temperature, and geochemical data tend to support this conceptual model (Hobba and others, 1979). Water is recharged rapidly into and through the friable Oriskany Sandstone and the cross-faulting serves as a drain, routing water to the Berkeley Springs (fig. 7).

Other structural offsets occur in addition to the fault trace depicted above. Structural offsets also occur at Rock Gap Run, North Fork Run, and Breakneck Run, which dissect Warm Springs Ridge and at Constant Run (Ziler Ford), which dissects Tonoloway Ridge at its southern edge. Cardwell and others (1968) mapped joints (approximately perpendicular to faults) coincident with the Constant Run, Rock Gap Run, and North Fork Run drainages. In addition to the Berkeley Springs, other regional springs occur in the study area.

Lessing and others (1991) report the relations between eight warm springs and the geology of eastern West Virginia by use of side-looking airborne-radar (SLAR). All of the springs discharge from Oriskany Sandstone or adjacent limestones of the Helderberg Group, although deep ground-water circulation to the Helderberg Group may be through the Oriskany Sandstone. They hypothesize that precipitation recharges the Oriskany Sandstone at ridgetop outcrops, then percolates deeply through the friable Oriskany Sandstone along fault zones where the ground water is geothermally heated. The ground water then flows through the highly permeable lineament, surfacing at the spring sites rapidly enough that the heat is not dissipated.

Ridge Hatchery Spring (site 93) surfaces in Breakneck Run, and Cacapon Spring (site 94) occurs in North Fork Run. Bedrock in lower Breakneck Run is highly fractured and occurs as blocky rubble along the mountainside and onto the valley floor. Zones of blocky fractured bedrock and bedrock rubble are very transmissive to water. Shallow subsurface flow can be heard beneath the rubblized deposits along the val-

ley floor in the Breakneck Run drainage. Recharge areas for the Ridge Hatchery Springs include Cacapon Mountain and potentially some local component of flow from Warm Springs Ridge. Ridge Hatchery Springs does not appear to have come from a deep flow system, as evidenced by water temperature and specific conductance.

Aquifer Properties

The specific capacity of a well is the rate of discharge of a well divided by the drawdown of the water level in the well and depends both upon the hydraulic properties of the aquifer and upon the construction and other features of the well. Specific-capacity estimates may be used to evaluate regional differences in transmissivity, which is the capability of an aquifer to transmit water through a unit width of aquifer under a unit hydraulic gradient (Heath, 1983). Short-term, single-well aquifer tests are used as the basis for estimating transmissivity of aquifer units in Morgan County. Transmissivity estimates are useful for developing a conceptual model of the aquifer and as input for modeling a ground-water flow system; however, transmissivity can be highly variable from one geographic location to another, both within a given geologic unit and among units. Additionally, an aquifer with a large transmissivity may be vulnerable to rapid contaminant transport, and aquifers with large transmissivities may be locally depleted if stressed beyond their sustainable yield.

Kozar and Mathes (2001) compiled aquifer characteristics from existing sources for the entire State of West Virginia. Two key characteristics discussed in that publication are specific capacity, defined as the rate of discharge of a well divided by the drawdown of the water level in the well; and transmissivity, defined as the rate at which water of a prevailing viscosity is transmitted through a unit width of aquifer under a unit hydraulic gradient (Kozar and Mathes, 2001). Although few of the data originated from wells in Morgan County, the values provide insight as to the range of values for hydraulic properties that occur throughout the extent of the hydrogeologic units. Specific-capacity and transmissivity data from throughout West Virginia (Kozar and Mathes, 2001) are summarized in table 1 for hydrogeologic units that occur in the study area. Except where otherwise noted, no data from the Kozar and Mathes (2001) report are available for the study area.

The largest specific capacity (600 (gal/min)/ft) and transmissivity (60,000 ft²/d) are in the Helderberg Group, Tonoloway Formation, and Wills Creek Shale. The Pocono Group and related MacCrady Formation, Hampshire Formation, and Chemung Group have the potential for moderate yields with maximum specific capacities of 8.67 (gal/min)/ft, 12.5 (gal/min)/ft, and 30.7 (gal/min)/ft, respectively, and maximum transmissivities of 2,300 ft²/d, 2,900 ft²/d, and 8,300 ft²/d, respectively. One site located in the study area has a specific capacity of 0.89 (gal/min)/ft and an estimated transmissivity of 240 ft²/d from the Hampshire Formation (Kozar and Mathes, 2001).

Table 1. Comparison of historical transmissivity and specific-capacity data (Kozar and Mathes, 2001) with results from this study for geologic units occurring in Morgan County, West Virginia.

Geologic unit	Historical transmissivity data (ft ² /d)			Transmissivity data from this study (ft ² /d)			Historical specific-capacity data [(gal/min)/ft]			Specific-capacity data from this study [(gal/min)/ft]		
	Minimum	Median	Maximum	Minimum	Median	Maximum	Minimum	Median	Maximum	Minimum	Median	Maximum
Pocono Group	33	210	2,300	15	15	15	0.17	0.83	8.67	0.13	0.13	0.13
Hampshire Formation	3	74	2,900	2	21	577	.01	.31	12.5	.06	.23	.86
Chemung Group	32	270	8,300	2	22.5	1,490	.12	1.35	30.7	.09	.46	3.75
Brallier Formation and Harrell Shale	3	72	390	2	4.5	96	.01	.28	1.60	.06	.27	1.00
Mahantango Formation	3	92	840	2	19	944	.01	.38	3.75	.08	.53	4.61
Marcellus Formation and Needmore Shale	16	170	1,300	54	128	327	.07	.86	7.00	.05	.40	.99
Oriskany Sandstone	19	82	1,500	(²)	(²)	(²)	.10	.35	8.00	(²)	(²)	(²)
Helderberg Group, Tonoloway Formation, and Wills Creek Shale	17	200	60,000	3	29.5	138	.07	.82	600	.06	.59	1.42
McKenzie Formation, Clinton Group, and Tuscarora Sandstone	6	23	1,100	2	124.5	266	.02	.11	6.50	.38	1.05	3.75

¹ One sample available.

² No samples taken.

The Mahantango Formation; Marcellus Formation and Needmore Shale; Oriskany Sandstone; and McKenzie Formation, Clinton Group, and Tuscarora Sandstone generally have much smaller maximum specific capacities of 3.75 (gal/min)/ft, 7.00 (gal/min)/ft, 8.00 (gal/min)/ft, and 6.50 (gal/min)/ft, respectively, and transmissivities of 840 ft²/d, 1,300 ft²/d, 1,500 ft²/d, and 1,100 ft²/d, respectively. The Brallier Formation and Harrell Shale have the smallest maximum specific capacity (1.60 (gal/min)/ft) and maximum transmissivity (390 ft²/d) (Kozar and Mathes, 2001) (table 1).

Single-well aquifer tests were done on private and public-supply wells to estimate transmissivity of aquifers in Morgan County (Appendix 1). Short-term aquifer tests, consisting of a drawdown phase and recovery phase, were done at most sites. In some instances, well production was so small and drawdown was so rapid that aquifer tests were discontinued prior to completion of the tests to avoid effects to the homeowner's well. In other instances, water levels were at a depth greater than 300 ft and could not be measured with available equipment. Some wells contained spacers or other equipment that precluded downhole placement of the pressure transducer necessary to conduct the test. Pump-cycling (the pump turning on and off) occurred at many of the sites.

During the drawdown test, the wells were pumped for approximately 30 minutes at a flow rate ranging from 0.80 to 60.0 gal/min. The analytical technique of Theis (1935) could not be used for tests in which interference occurred because of pump cycling. Transmissivity was computed only from recovery data at those sites.

When the pumping is terminated, water from within the aquifer flows toward the well bore and, over time, the water level recovers to static conditions. During this aquifer recovery period, water-level data typically were recorded for approximately 60 minutes, with some recorded for as long as 131 minutes. Recovery data were analyzed by using the techniques of Theis (1935), which are based upon certain simplifying assumptions of aquifer properties: a homogeneous (identical properties throughout), isotropic (same properties with the same values when measured along axes in all directions), confined aquifer that is uniform in thickness and infinite in extent. The well is assumed to be infinitesimally small in diameter and to fully penetrate a confined aquifer. These are considered ideal aquifer conditions. Conditions in Morgan County are nonideal conditions; however, the technique was applied to generate a range of values for the aquifer properties of hydrogeologic units in the study area.

With the exception of the highly friable zones of the Oriskany Sandstone, aquifers in Morgan County are very heterogeneous and anisotropic. Typically, well-completion reports for the county indicate the presence of only one to two isolated zones of production along the length of the wellbore. Parameter estimates in poorly transmissive zones of the aquifer are subject to error because of unaccounted-for wellbore storage effects.

Tiedeman and Hsieh (2001) tested fractured crystalline-rock aquifers and concluded that the borehole connects highly permeable zones separated by impermeable rock. Hsieh and Shapiro (1996) further state that single-well tests show the effects of only the few highly transmissive fractures and are not representative of the overall bulk-rock properties. Schad and Teutsch (1994) suggest that the Theis solution can be used for comparison of values within a single dataset rather than representing the values as valid for a rock type. For the purposes of this report, the transmissivity values are used to evaluate the variability within aquifers and among rock types in Morgan County and should not be used for direct comparison with data from other areas, although they may provide a basis for design of more complex aquifer tests.

Parameter estimates from drawdown and recovery tests agreed to within 174 percent of each other. Parameter estimates from recovery tests exceeded those from drawdown tests at 74 percent of the sites. Reported values reflect the average transmissivity derived from both the pumping and recovery tests. The median transmissivity for the entire dataset (91 wells with 21 pumping and 70 recovery tests) was 23.0 ft²/d, with a minimum of 2.0 ft²/d and a maximum of 1,490 ft²/d (Appendix 1).

Linear regression analysis of Morgan County well sites shows no significant correlation ($r^2 = 0.00$) between well depth and transmissivity. The median transmissivity for wells with depths less than or equal to 150 ft below land surface is 27.2 ft²/d. The median transmissivity for wells greater than 150 ft below land surface is 21.4 ft²/d. This transmissivity distribution indicates that discrete flow zones, although smaller at depth, are not restricted to the upper portion of the aquifers. The data also are consistent with data from nearby Berkeley County, W. Va., that indicate a slight reduction in aquifer yield below a depth of approximately 150 ft below land surface (Shultz and others, 1995). Transmissivities range over more than two orders of magnitude. The largest calculated transmissivities were at site 77, which lies approximately at the contact between the Chemung Group and the Brallier Formation with a transmissivity of 1,420 ft²/d, and at site 85, which lies approximately at the contact between the Mahantango and Brallier Formations, with a transmissivity of 944 ft²/d (fig. 8). In the study area, larger aquifer transmissivity typically occurs at or near geologic contacts and in zones of cross-faulting or thrust-faulting.

Transmissivity data for geologic units located in the study area from previous studies are compared with those from this study (table 1). In general, the transmissivities from this study were smaller than those previously documented for the Hampshire Formation, Chemung Group, and Brallier, Mahantango, Marcellus, Tonoloway, and Clinton Formations (Kozar and Mathes, 2001). Only one site was sampled in the Pocono Group, and that transmissivity was nearly an order of magnitude smaller than those reported by Kozar and Mathes (2001).

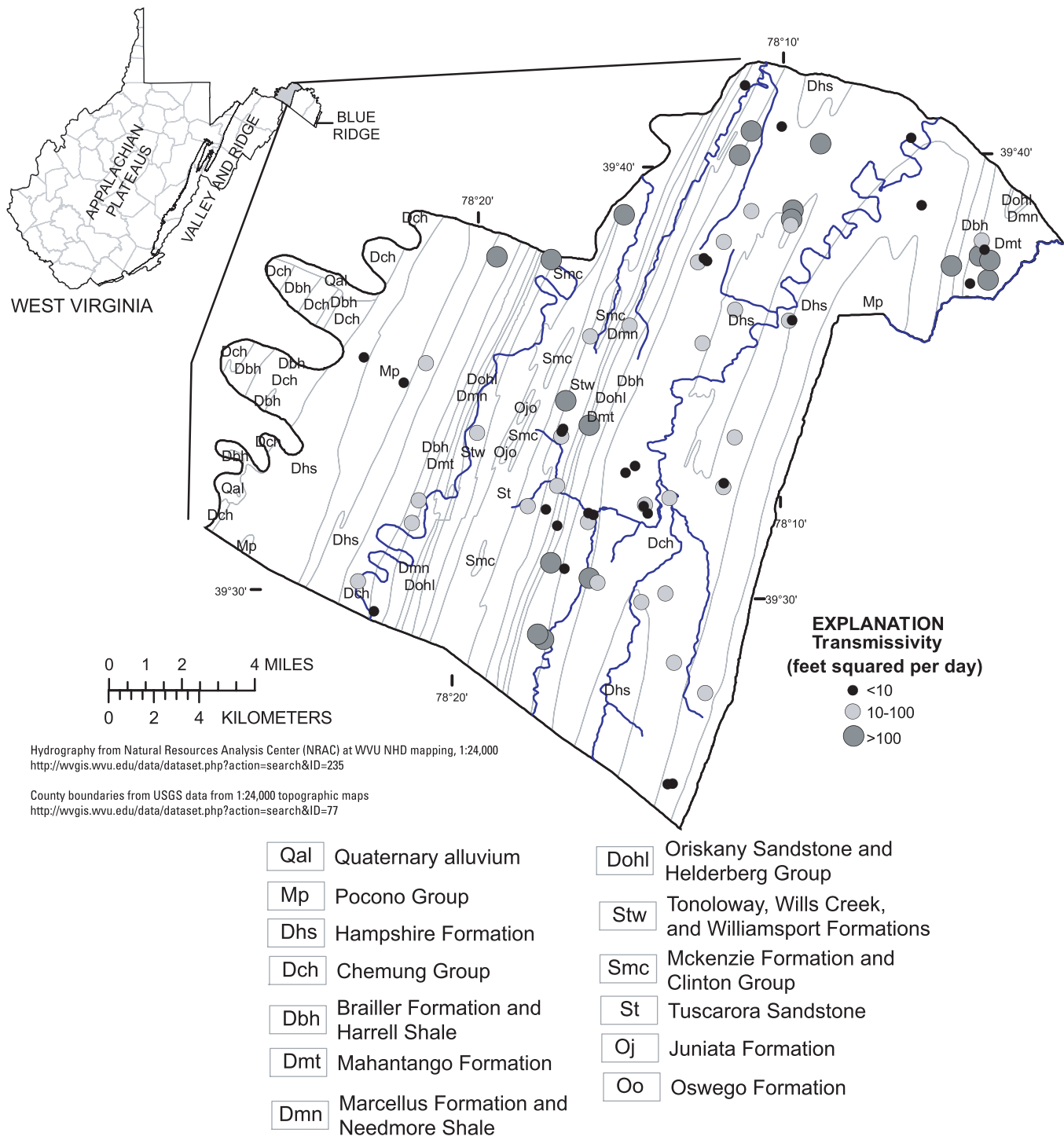


Figure 8. Locations and magnitudes of transmissivities, Morgan County, West Virginia. Geology after Cardwell (1968).

Values for previous transmissivity data within the study area for the Marcellus and Hampshire Formations (Kozar and Mathes, 2001) are within the range of values obtained during this study. Site-specific hydrogeologic details are difficult to assess from the literature; however, depths for the sites with the largest transmissivities reported by Kozar and Mathes (2001) tend to be shallower than the typical well depth in Morgan County. Well yields decrease appreciably at depths below approximately 150 ft below land surface because of a decrease in fracturing (Shultz and others, 1995). Lithostatic pressure from the mass of overlying rock tends to compress fracture partings. Transmissivity of the Oriskany Sandstone was not sampled during this study because it has narrow outcrops along Warm Springs Ridge where few dwellings exist, thus limiting availability of sampling sites.

Site-specific geographic location cannot be used to predict transmissivity. Wells in close geographic proximity within the study area have transmissivities that vary by more than two orders of magnitude. The exception is that aquifers tested along ridgetops tend to have smaller transmissivities than those in valley settings, and zones of production tend to be deep, in excess of 400 ft below land surface.

Water Levels

Ground-water flow paths are a reflection of the complex geologic setting in the study area. Static water-level data were obtained at 84 sites throughout Morgan County (fig. 9).

Data density is not adequate to predict local or site-specific flow systems with any certainty or to develop countywide static water-level contours. These data are used to develop a conceptual model of the generalized ground-water-flow patterns, which will be addressed progressively from west to east in the study area.

The Sideling Hill Syncline is a recharge area with recharge draining to depth through the fractured Hampshire Formation. Wells range from 380 to 460 ft deep and have low yields from fracture sets at depth. Static water levels range from approximately 100 ft to greater than 220 ft below land surface. Bare Ridge, Road Ridge, and Tonoloway Ridge may form structural barriers to the east. Ground-water flow appears to be northward within the Willett Run drainage, as the static water level is several hundred feet lower at the north end of that drainage than at its southern reaches.

Water levels at locations near the Cacapon River range from 20 ft to approximately 60 ft below land surface. The ground-water gradient indicates that ground-water flow follows a path similar to the Cacapon River from the southern boundary of the study area, through a water gap at Constant Run that occurs at the south end of Tonoloway Ridge, and then northward. The Constant Run water gap is coincidental with cross-faulting that occurs in that gap between the south end of Tonoloway Ridge and the north end of Little Mountain (fig. 3). That fault trace appears to be present at the Potomac River to the west and is evident as far east as Cacapon Mountain.

It appears to act as a subsurface drain to the north-northeast, along the Cacapon River Fault. There may be a ground-water divide north of the fault trace, but additional data are required to validate that hypothesis.

Cacapon Mountain is a recharge area. The hinge line of the Cacapon Mountain Anticline may be a drainage divide with ground water flowing to the west and the east; however, the Boundary Fault underlies Cacapon Mountain. This slip-fault extends as deep as the Martinsburg Decollement, several thousand feet below land surface and may serve as a bounding fault to the west or the fault zone may serve as a conduit for deep flow to the ground-water system east of the anticline. Grimsley (1916) characterizes this area as having a series of crumpled, fluted folds mapped at the surface, which tends to support the potential for flow through fractures or joints. Alternatively, ground water could flow toward the west and may leave the system through the Cacapon River Fault zone (fig. 4). Deep flow to the east from the crest of Cacapon Mountain could account for head needed to produce the artesian conditions (water levels within 10 feet or less from land surface) that occur in Sir Johns Run and east of Warm Springs Ridge (sites 33, 34, 61, 70, 73, and 74) (fig. 2).

Hydrogeologic data are not available along the crest of Cacapon Mountain (the anticline hinge-line), but one site (59) near the top of the mountain has a static water level of approximately 1,080 ft above sea level, considerably higher than others in the Sir Johns Run drainage. Sir Johns Run drainage is described by Grimsley (1916) as a Silurian valley lying between the resistant sandstone ridge that caps Cacapon Mountain and Warm Springs Ridge also formed by another resistant sandstone unit, the Oriskany Sandstone. The drainage between was formed by erosion of intervening, less resistant shales. Ground-water level altitudes in Sir Johns Run drainage tend to be lower toward the north, from approximately 890 ft to 710 ft at the north end of the drainage, indicating that ground water flows northward. There is an anomaly in the water levels in Warm Springs Run drainage north of the town of Bath. This localized perturbation of the flow path may be the result of dewatering operations at a mine in Warm Springs Ridge.

The Third Hill Mountain complex lies along the hinge line of the Meadow Branch Synclinorium. Localized recharge occurs along this ridge, which also acts as a ground-water divide with flow to the northeast and northwest following the north-south structural trends with the units plunging slightly to the north. Water levels confirm this conceptual model. Six wells were hydraulically tested in the area east of Third Hill Mountain; the transmissivity measured at four of those wells was more than 100 ft²/d, whereas the transmissivity at two of the sites was appreciably less than 20 ft²/d. This variability may be complicated by the occurrence of the Woods thrust fault in that area that offsets already complex geology. Grimsley (1916) reported the presence of striated and polished rock surfaces, known as slickensides, on the faults near Third Hill Mountain. Slickenside rocks may act as barriers to flow rather than as conduits typical of extensively fractured rock.

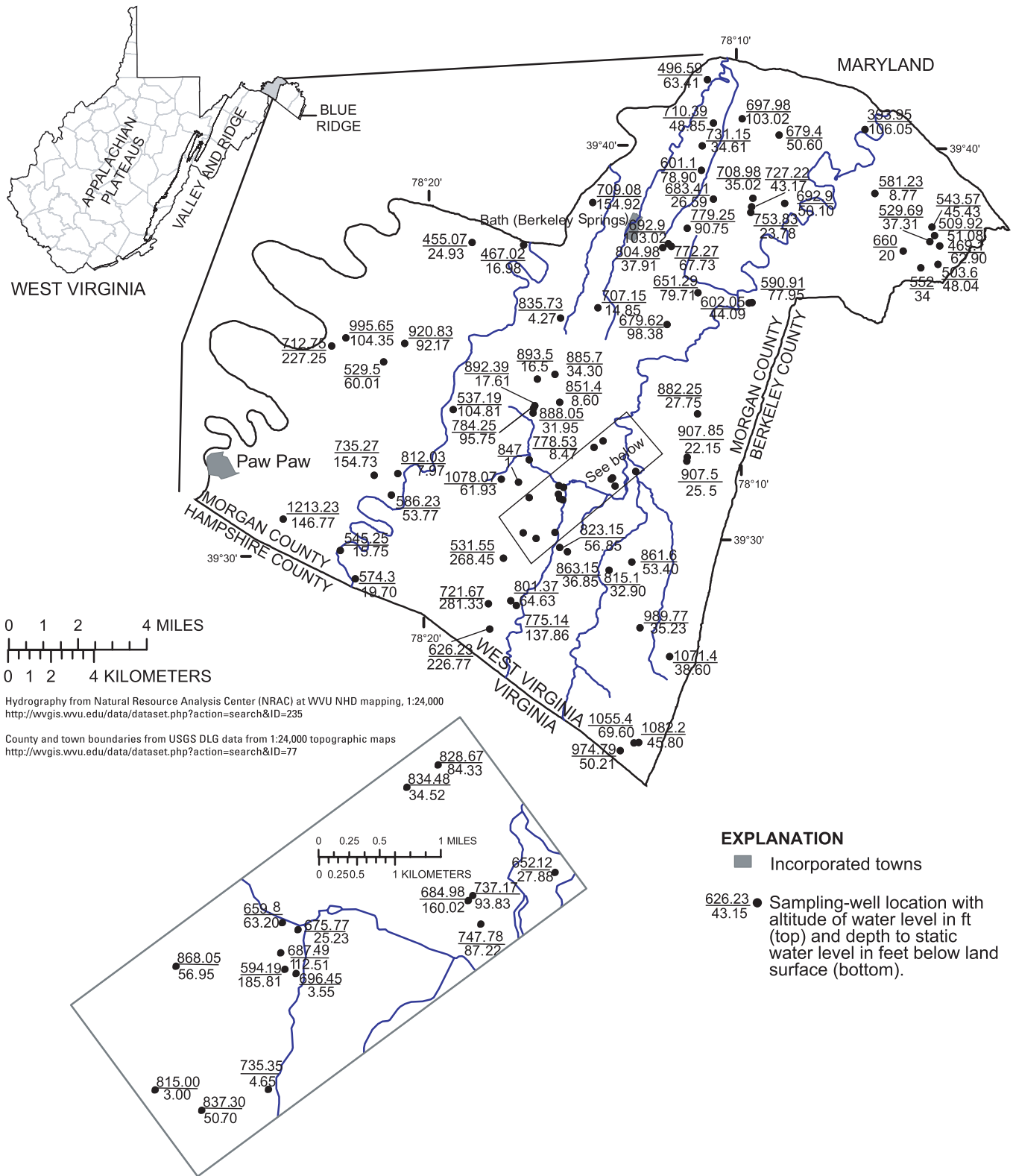


Figure 9. Ground-water-level altitudes and depth to static water levels in Morgan County, West Virginia.

Aquifer Geochemistry and Ground-Water Quality

Ninety-one wells and eight springs in the study area were evaluated for aquifer geochemistry and ground-water-quality characteristics (Appendixes 2 and 3). Geochemical properties and ground-water quality are largely affected by composition of the rocks in the aquifer. Additionally, contaminants may be introduced by human activities or other environmental factors. USEPA has established regulatory water-quality standards for specified constituents for the point of delivery from public water supply systems. The standards are based upon potential human-health or esthetic effects (U.S. Environmental Protection Agency, 2005a, b, c).

Geochemical Properties

Geochemical properties of aquifer systems are naturally variable and dependent upon the mineralogic composition of the rocks encountered along the ground-water flow path. The three major sedimentary rock types (carbonates, sandstones, shales) all occur in Morgan County. Limestone and dolomite rocks have a predominant calcium-bicarbonate or calcium-magnesium-bicarbonate geochemical signature, respectively. Pure sandstone is quartzite that may be cemented by silica or by calcium carbonate or may contain little cementation. In Morgan County, cementation of the Oriskany Sandstone varies depending upon location. Sandstone cemented by silica may have up to 99 percent pure silica content, which will be reflected in the geochemical signature. Shales primarily consist of various clay minerals and quartz. Shales formed under anaerobic conditions may retain organic matter and may have a large sulfur content. Other shales may be enriched in metals such as arsenic, barium, boron, chromium, copper, fluoride, gallium, lithium, mercury, selenium, titanium, uranium, yttrium, vanadium, and zinc (Brownlow, 1979). Ground water may reflect a mixture of the geochemical characteristics as it flows progressively through one rock unit and then through another. Complex geochemical interactions may result in dissolution of certain constituents and precipitation of other constituents over time. Geochemical properties and concentrations of water-quality constituents for wells and springs sampled during this study are presented in Appendixes 2 and 3, respectively.

Water-Quality Characteristics

Where appropriate, concentrations or other water-quality values are compared with U.S. Environmental Protection Agency (USEPA) Maximum Contaminant Levels (MCLs) and Secondary Maximum Contaminant Levels (SMCLs) drinking-water standards established for water at the point of delivery from public drinking-water supply systems. The MCLs are

health-based drinking-water standards, whereas SMCLs are established on the basis of cosmetic effects such as skin or tooth discoloration, esthetic effects such as taste, odor, or color, and potential damage to plumbing systems such as corrosiveness (U.S. Environmental Protection Agency, 2005b and d). No water-quality standards are established for water from private wells or springs.

pH

The USEPA SMCL states that pH should range from 6.5 to 8.5 (U.S. Environmental Protection Agency, 2005d). A pH of 7.0 indicates a neutral aqueous solution. A pH higher than 7.0 is indicative of alkaline water, and a pH lower than 7.0 is indicative of acidic water. The corrosiveness of water to pipes and plumbing increases as the pH decreases. Excessive alkalinity in water also can attack metals. The pH of water at 32 sites was lower than 6.5 pH units. No samples had pH in excess of 8.5 pH units (fig. 10). The maximum pH measured at any of the sites was 7.7 pH units. The median pH of samples from the Clinton Formation was 5.55 and from the Marcellus Formation was 5.75. The pH in all other units combined was 6.75. The pH of the Marcellus and Clinton Formations is appreciably more acidic than the other units, which tend to be more neutral.

Geochemically, pH is a significant control on the dissolution of certain naturally occurring metals. Alternatively, pH is controlled by the aquifer material. For instance, in oxidizing conditions, pH may be low because of dissolution of ferrous-containing minerals, whereas pH may rise because of hydrolysis of silicates resulting in a geochemical environment that is both alkaline and reducing (Garrels and Christ, 1965). In the study area, alkalinity increases at pH measurements greater than approximately 6 (fig. 11). Concentrations of silica increase appreciably at pH measurements greater than 6 (fig. 12).

Specific Conductance

Specific conductance is a measure of the ability of a substance to conduct an electrical current and is reported in units of microsiemens per centimeter at 25 degrees Celsius ($\mu\text{S}/\text{cm}$). Specific conductance of water increases as the concentrations of dissolved ions increase. More mineralization produces a higher specific conductance. Specific-conductance values from the Tonoloway Formation had the largest variability among the hydrogeologic units sampled, ranging from 318 to 737 $\mu\text{S}/\text{cm}$, with a median value of 445 $\mu\text{S}/\text{cm}$ and a standard deviation of 118.7 $\mu\text{S}/\text{cm}$ (fig. 13). Specific-conductance values from the Brallier Formation had the least variability, ranging from 202 to 322 $\mu\text{S}/\text{cm}$, with a median value of 270 $\mu\text{S}/\text{cm}$ and a standard deviation of 35.5 $\mu\text{S}/\text{cm}$. Samples from all other formations varied by up to an order of magnitude from the minimum to the maximum specific-conductance value.

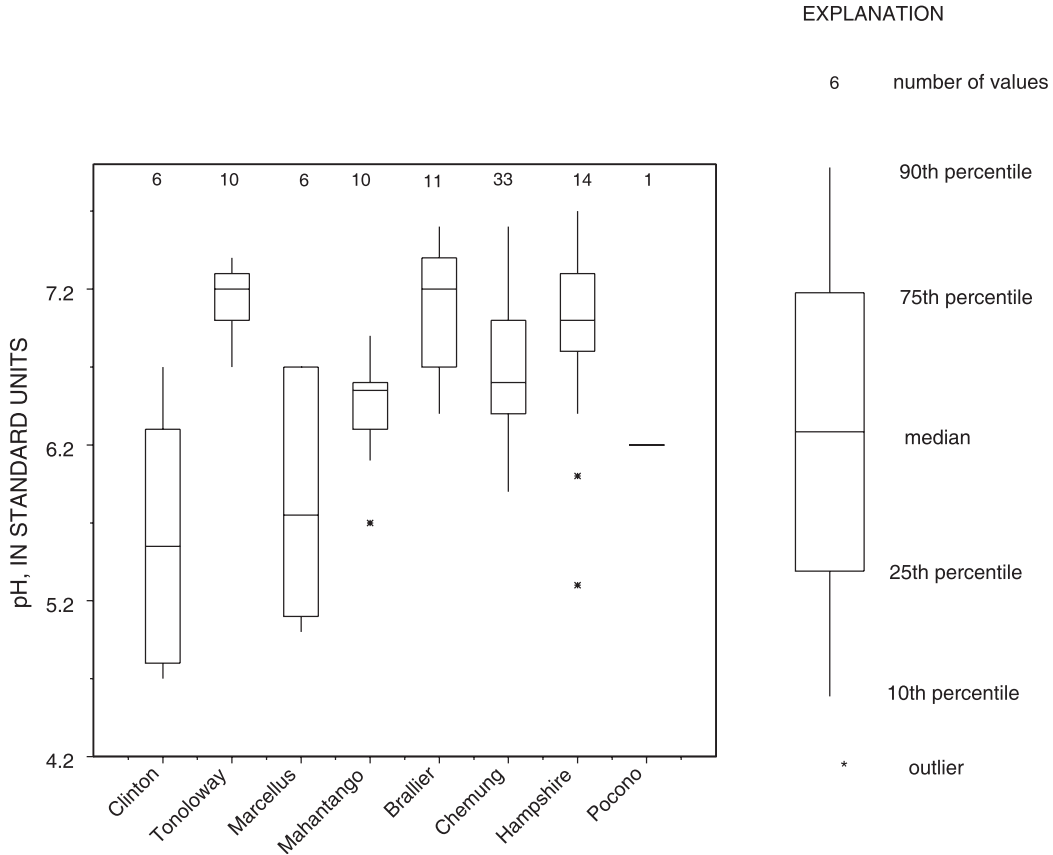


Figure 10. Distribution of pH values in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.

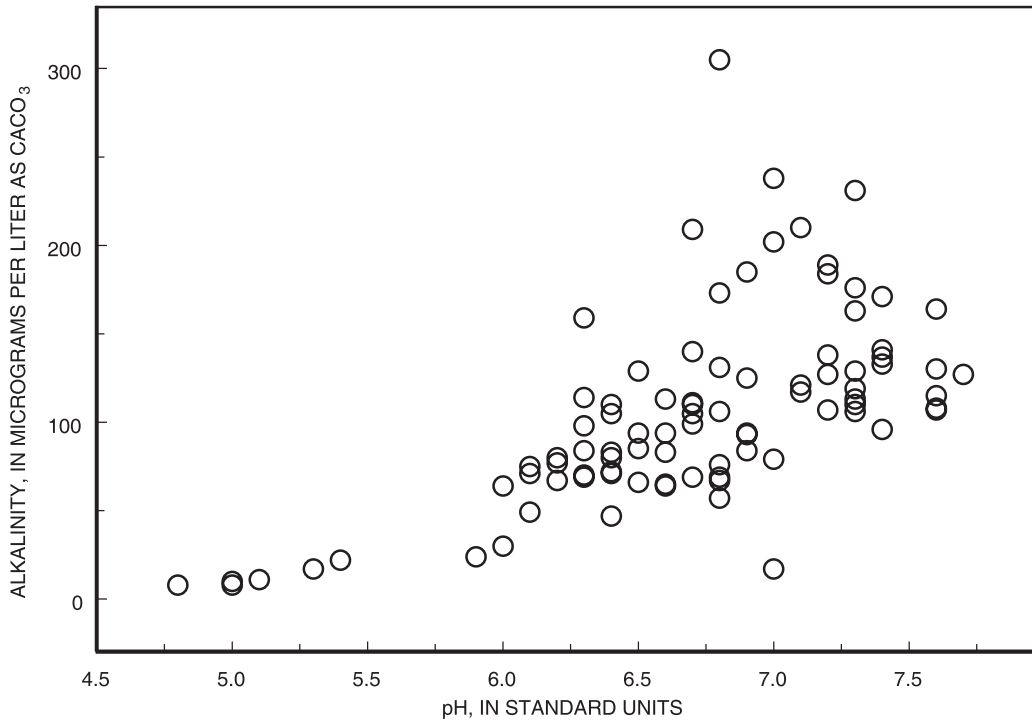


Figure 11. Relation between concentration of pH and alkalinity in ground-water samples from Morgan County, West Virginia.

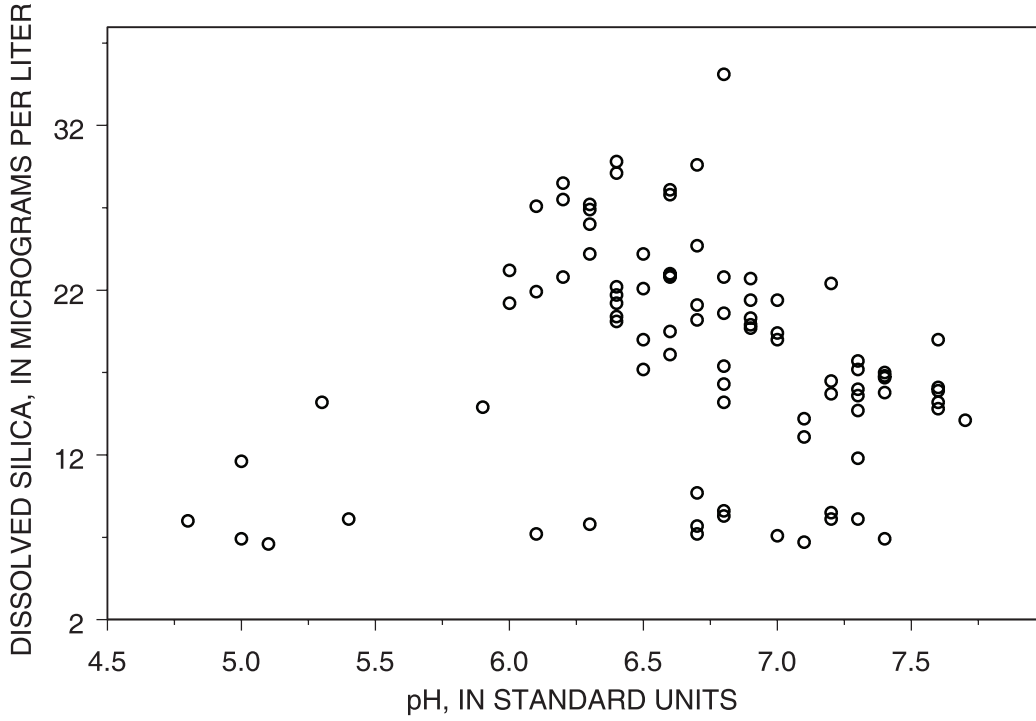


Figure 12. Relation between concentration of pH and dissolved silica in ground-water samples from Morgan County, West Virginia.

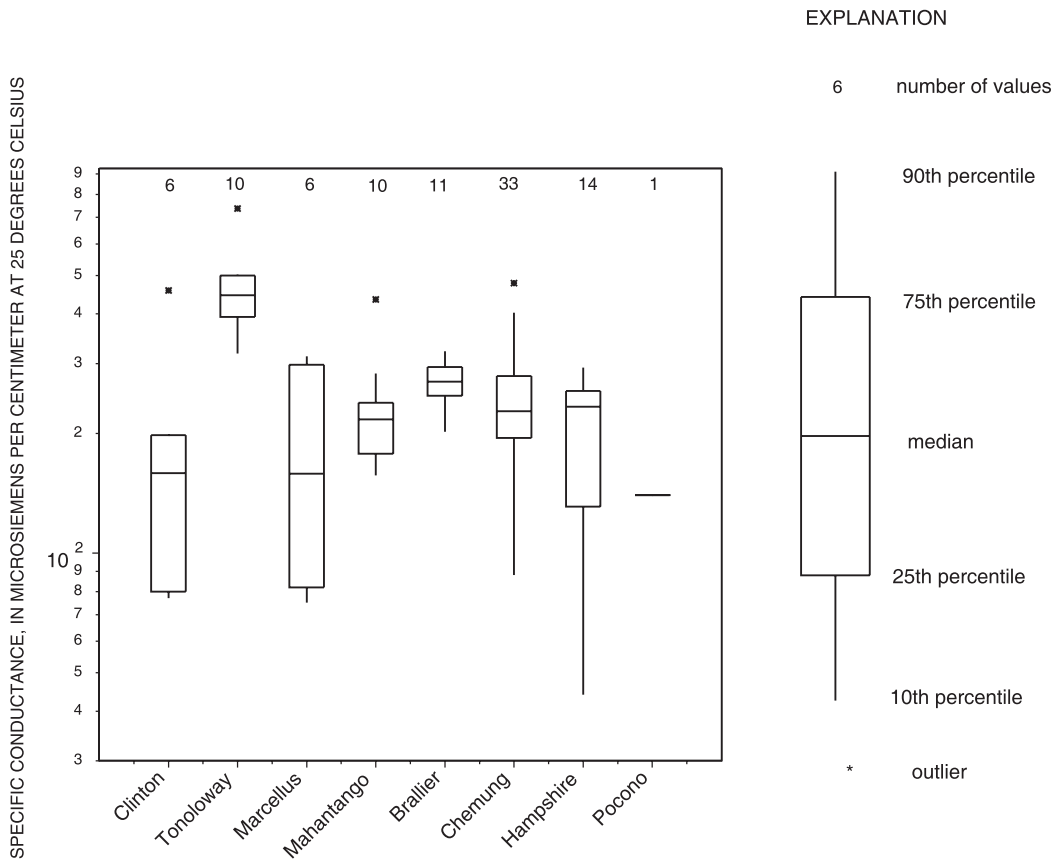


Figure 13. Distribution of specific conductance in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.

Temperature

Ground-water temperature responds to heat flow from the Earth's interior and, near the Earth's surface, seasonally variable heat from the sun (Heath, 1983). The near-surface, seasonally variable zone is approximately 30 to 80 ft thick, and ground-water temperatures in that zone may vary seasonally by as much as 5 to 10°C (Heath, 1983). Mean annual ground-water temperature of this shallow zone is estimated to be in the range of 10 to 13°C in the study area (Heath, 1983). At depths below 100 ft below land surface, ordinary geothermal temperature gradients range from approximately 0.6 to 0.7°C per 100 ft (Driscoll, 1986). Elevated ground-water temperatures may indicate that the ground-water flow path is through a deep aquifer system in which geothermal heating has taken place. Meteoric water that is recharged at relatively high elevation on the crest or limb of an anticline may flow along bedding planes or fracture zones deep into an adjacent syncline where geothermal heating takes place. The flow then rises to or near

the surface, where the same bed crops out in another anticline (Reeves, 1932) or through zones of weakness in the bedrock such as those caused by transverse or thrust faults. Perry and others (1979) evaluated geothermal heat flow in Bath County, Va., which lies in a physiographic setting similar to that found in Morgan County, W. Va. In that study, the Brallier Formation of Devonian age was found to have a geothermal gradient of approximately 0.4°C per 100 ft.

The ground-water temperature was measured at all well and spring sites in the study area. The temperatures measured at the well sites ranged from 10.9 to 16.0°C with a median temperature of 13.0°C (fig. 14). These temperatures indicate that ground water in the study area is derived from flow systems potentially ranging from a near-surface system, less than 200 ft below land surface, to a system as deep as 1,500 ft or more. However, depending upon the flow path, geothermally heated ground water may cool within the flow system before its interception by a well or its issuance at the surface.

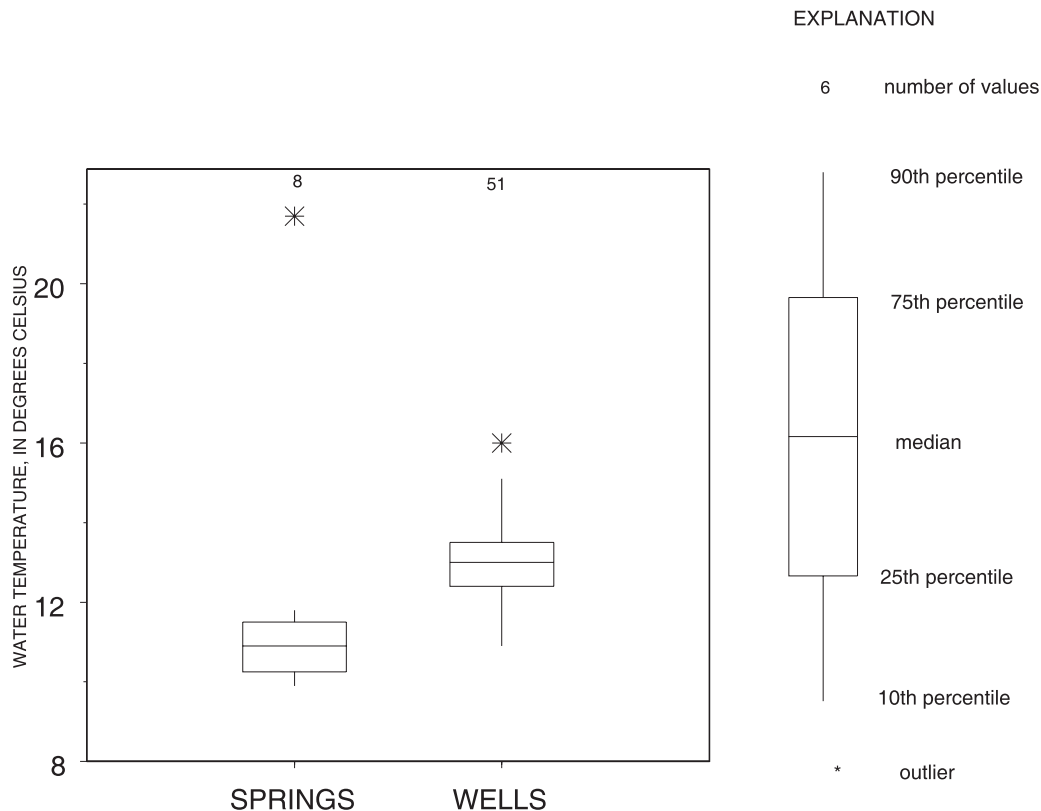


Figure 14. Distribution of temperature in ground-water samples from springs and wells located in Morgan County, West Virginia.

Temperatures at the spring sites ranged from 9.9 to 21.7°C. The median temperature of the springs was 10.9°C. Lessing and others (1991) reported temperatures from geothermally heated warm springs in Morgan County ranging from 15.3°C at Hancock Spring to 22.2°C at Berkeley Springs. Hancock Spring was not sampled during the current study. The Berkeley Springs discharge from five major orifices, one of which is the Ladies Spring. The Ladies Spring (site 100) sample was 21.7°C and was the only spring sampled during this study that had a temperature indicative of a deep flow system. Water produced from the Marcellus Formation was generally warmer than the other formations, ranging from a minimum of 12.3°C to a maximum of 16.0°C and with a median of 13.8°C (fig. 15). Water produced from the Brailler Formation had a similar temperature profile, ranging from a minimum of 12.4°C to a maximum of 14.2°C and with a median of 13.9°C. Water samples produced from the Mahantango Formation had the smallest temperature range from 11.6 to 13.0°C and a median of 12.4°C. The results indicate that water sampled from the Marcellus and Brailler Formations may flow along deeper flow paths than those in the Mahantango Formation.

Alkalinity

Alkalinity is defined as the capacity of a solution to neutralize acid at a designated end-point. Alkalinity may be produced by noncarbonate anions such as sulfate, phosphate, and nitrate. However, in most natural water, nearly all alkalinity is produced by dissolved carbonate and bicarbonate and is commonly expressed in terms of equivalent calcium carbonate or carbonate alkalinity (Hem, 1985). Precipitation contains atmospheric carbon dioxide and dissolves additional carbon dioxide as it infiltrates through the soil zone. Carbonate rocks, such as limestone and dolomite, are composed primarily of calcium carbonate (CaCO₃). As ground water containing carbon dioxide comes in contact with carbonate rocks, the calcium carbonate in the rocks is dissolved, producing bicarbonate. The USEPA (2005b) has not established a water-quality standard for alkalinity. Concentrations of alkalinity in samples from the Tonoloway Formation were considerably larger than from any other aquifer. The median concentration of alkalinity for the Tonoloway Formation was 184 mg/L, the minimum was 69 mg/L, and the maximum was 305 mg/L. The median

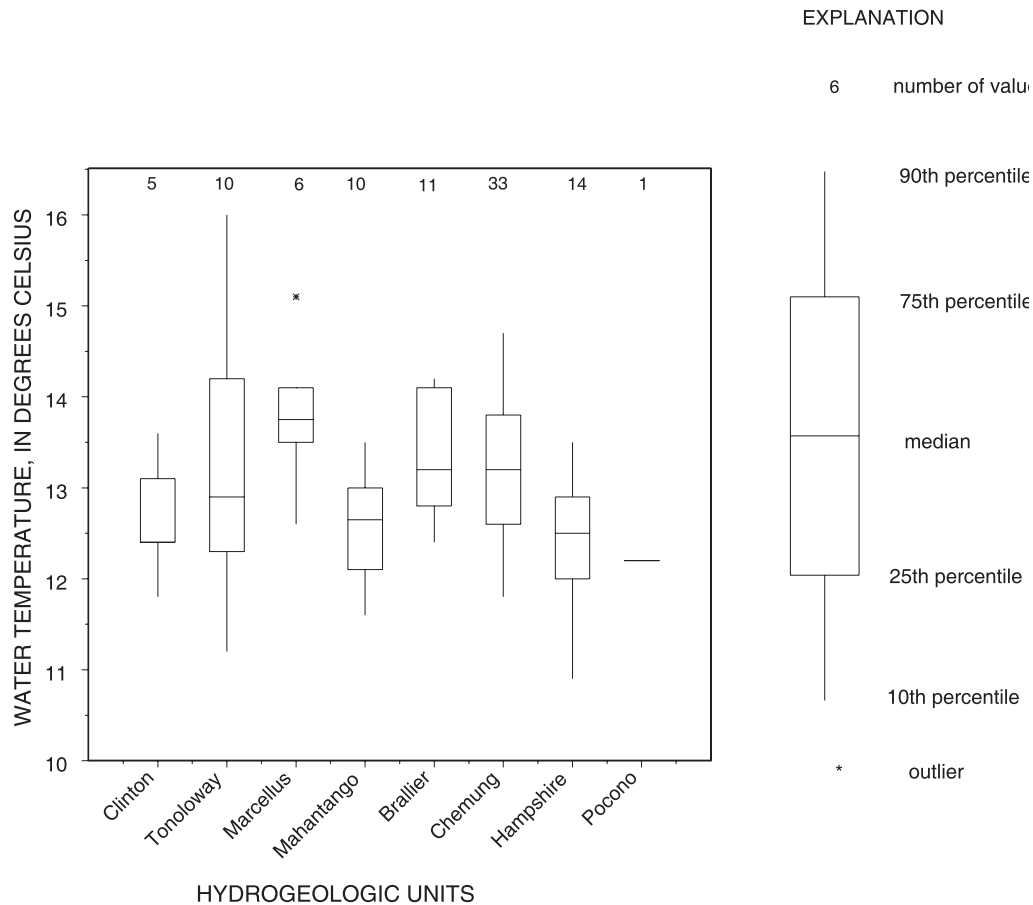


Figure 15. Distribution of temperature in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.

concentration of alkalinity for all other ground-water samples was 94 mg/L, the minimum concentration was 8 mg/L, and the maximum was 238 mg/L. The concentration of alkalinity of all of the ground-water samples collected was produced primarily by bicarbonate.

Water Hardness

Water hardness is caused primarily by calcium and magnesium ions in solution. In this report, the convention used for expressing hardness is “hardness as CaCO_3 ” and is calculated as described by (Hem, 1985). The USEPA (2005b) has not established a standard for hardness, which is typically associated with effects observed in the use of soap or with the encrustations left by some types of water when they are heated (Hem, 1985). Hardness is commonly classified as follows: soft (0 to 60 mg/L); moderately hard (61 to 120 mg/L); hard (121 to 180 mg/L); and very hard (greater than 180 mg/L) (Durfor and Becker, 1964). On the basis of those criteria, approximately 13 percent of the samples analyzed were soft water, 56 percent were moderately hard water, 18 percent were hard water, and 13 percent were very hard water. Ground-water hardness as CaCO_3 in the Tonoloway Formation ranged from 80 to 400 mg/L with a median concentration of 220 mg/L, considerably larger than from other aquifers. Concentrations of hardness as CaCO_3 from all other formations ranged from 16 to 260 mg/L with a median concentration of 98 mg/L.

Dissolved Oxygen

Oxygen is supplied to ground water through recharge and by movement of air through the unsaturated zone. Recharge water can be expected to contain concentrations of DO similar to surface water. The equilibrium concentration of DO in water is primarily a function of water temperature and atmospheric pressure. Traditional theory holds that water in deep ground-water systems is typically undersaturated with respect to oxygen. As ground water flows through an aquifer system, DO reacts with available, oxidizable, aquifer material such as pyrite (FeS_2) and siderite (FeCO_3) or naturally occurring organic material resulting in a decreased concentration of DO. In settings with shallow soil overlying permeable fractured rock, DO at detectable concentrations may persist throughout the flow system (Freeze and Cherry, 1979). Concentrations of DO from 2 to 8 mg/L have been measured in ground water from deep aquifers in the south-central Great Basin Physiographic Province (Fenneman, 1946). The residence time for ground water in these aquifers ranges from thousands to more than 10,000 years (Winograd and Robertson, 1982). Edmunds and others (1982) reported appreciable concentrations of DO in recently recharged water in a sandstone aquifer in England, whereas downdip concentrations of DO in the same aquifer were depleted to less than 0.1 mg/L. Hydrogeochemical calculations resulted in an estimate of an approximate 5,000-year

residence time for the water in that aquifer to have evolved to those low concentrations.

For this study, concentrations of DO ranged from estimated values of less than 0.1 to 10.3 mg/L. The 10.3 mg/L concentration represents 100 percent saturation of DO at the measured temperature and barometric pressure. Water was observed to be cascading down the wellbore at sites 30, 39, 40, and 83, and air bubbles became entrained within the homeowner’s plumbing system at site 47; therefore, concentrations of DO at those locations are assumed to be elevated above that which would have occurred naturally within the aquifer. The concentration of DO at several other sites was elevated for no observed reason. It is possible that ground water flows so rapidly through the fracture systems that geochemical reduction of the concentration of DO is minimized. It also is possible that turbulent flow occurs within open fracture systems or solutionally enhanced conduits, thus entraining oxygen in place. Pumping with rapid drawdown in low-production domestic wells may induce mechanical entrainment of DO.

Turbidity

Turbidity is a measure of the cloudiness of water (U.S. Environmental Protection Agency, 2005b). Turbidity occurs naturally in ground water, depending upon the aquifer medium. Wells may produce nonturbid water at low pumping rates, but at higher pumping rates, turbidity may increase dramatically. The increased turbidity could be from turbid recharging water being pulled into the stressed aquifer or from turbulent flow that picks up sediment in fractures or conduits within the aquifer and suspends it in the ground water (Hobba and others, 1972). The maximum turbidity measured during this study was 330 nephelometric turbidity units (NTU), the median turbidity was 3.9 NTU, and 22 samples had turbidity of less than 1 NTU. USEPA has established a maximum for turbidity in public supply systems as 5 NTU prior to treatment, with more stringent requirements for public-water-supply systems that filter (U.S. Environmental Protection Agency, 2005b). USEPA requirements are established on the basis of turbidity caused by disease-causing organisms typically found in surface water, not ground water. Thirty-five of the wells sampled had turbidity measurements greater than 5 NTU at the pumping rates sustained over approximately a 30-minute time interval. Turbidity measurements typically did not stabilize, but increased throughout the pumping test, indicating that the aquifer could not sustain production at the indicated pumping rate without inducing turbidity.

Inorganic Constituents

Inorganic constituents are those that do not contain carbon. Ninety-one wells and eight springs were sampled in the study area for a variety of inorganic constituents; results are summarized in Appendixes 2 and 3, respectively.

Major Ions

Major inorganic ions are calcium, magnesium, sodium, potassium, bicarbonate, chloride, and sulfate, which typically occur in natural water in concentrations of 1 mg/L or larger. These constituents exist in pairs of cations and anions, which are indicative of the mineralogy of the hydrogeologic setting through which the water has flowed. For instance, calcium-bicarbonate-dominant ground water is indicative of a limestone aquifer, calcium-magnesium-bicarbonate water may be indicative of a dolomite aquifer, and sodium-chloride dominant ground water is typical of a sedimentary setting rich in evaporite salts. Geochemical characterization becomes more complex where the aquifer system consists of a mixture of rock types or if the ground water has flowed through differing types of rock units.

Ground water within a study area may be classified by use of a trilinear diagram suggested by Piper (Hem, 1985). To use this diagram, the concentration of each major cation and anion is expressed as a percentage of the total milliequivalents per liter in each sample. The percentage of major cations and anions in each sample is then plotted as a single point on the lower left and lower right triangles of the diagram, respectively. These points are projected to a single point on the plotting field, which lies between the two triangles. For example, if the point falls near the left point of the plotting field, then the water may be classified as a calcium-bicarbonate-type water. If a point falls near the lower right tip of the plotting field, the water is classified as a sodium-potassium-bicarbonate-type water. The trilinear method describes the percentage of the major cations and anions. That is, the percentage of each constituent or group of constituents is represented by where a point lies in the plotting field; however, this graphical method does not indicate the concentration of the dissolved constituents. Geochemical signatures of ground-water and spring samples analyzed for this study are shown in figure 16.

The predominant geochemical signature of ground water and spring water in Morgan County was calcium-bicarbonate-type water; however, there also were vastly differing ground-water samples with predominantly sodium-bicarbonate and predominantly calcium-magnesium-sulfate signatures.

Most of the spring-water samples were of the calcium-bicarbonate type (fig. 16). A distinct exception was that of spring site 94 at Cacapon State Park, which showed a calcium-sulfate-chloride signature similar to that of the Marcellus Formation. Spring site 93 has a geochemical signature that lies between the calcium-bicarbonate and calcium-sulfate-chloride signatures, indicating that the ground water has flowed through some combination of limestone and shale units and plots close to ground-water samples from the Hampshire Formation (fig. 16).

The Pocono Group, Chemung Group, and Hampshire and Tonoloway Formations produce predominantly calcium-bicarbonate waters. Some samples had an elevated sulfate signature. Hobba and others (1972) attributed the possible source of an increased sulfate signature in the Tonoloway Formation as the presence of anhydrite (CaSO_4) and salt (NaCl) deposits. Anhydrite and salt are known to occur at depth west of the study area but generally are thought to have been removed by ground-water circulation in the Potomac River Basin; however, residual deposits that have not been completely dissolved remain in rocks in the study area. The Hampshire Formation and Chemung Group also had some samples that had an increased sodium-bicarbonate signature. One sample (site 8) that is mapped at the surface in the Hampshire Formation had a nearly 100 percent sodium-bicarbonate composition, a very soft water. A narrow section of the Onesquethaw Group occurs at the surface near that well location. Deposits of glauconite, which is used as water softener, are known to occur in that group. Thicknesses of the Hampshire Formation and underlying units are not well known in that area, which is structurally complex. It is possible that the 300 ft well at that site is tapping water that originates from the Onesquethaw Group. Most samples from the Clinton Formation had a calcium-bicarbonate signature, with one exception: a calcium-sodium sulfate water that bore a signature similar to that from a warm spring in the Oriskany Sandstone (Hobba and others, 1972).

Bromide and Fluoride

In addition to chloride, analyses were done for two other halides: bromide and fluoride. USEPA (2005b and d) has not established a drinking-water standard currently (2006) for bromide. The largest concentration of bromide detected was 0.13 mg/L at one site. USEPA (2005b and d) has established standards for fluoride. An MCL of 4.0 mg/L is established on the basis of potential health effects, including bone disease or mottled teeth. The SMCL of 2.0 mg/L is established for naturally occurring fluoride that may cause tooth discoloration or pitting during the formative period prior to eruption of teeth in children (U.S. Environmental Protection Agency, 2005d). None of the spring or well samples had concentrations of fluoride greater than the MCL or the SMCL. Approximately 42 percent of the samples had concentrations lower than the detection limit of 0.2 mg/L. The largest concentration was 0.3 mg/L, at three sites.

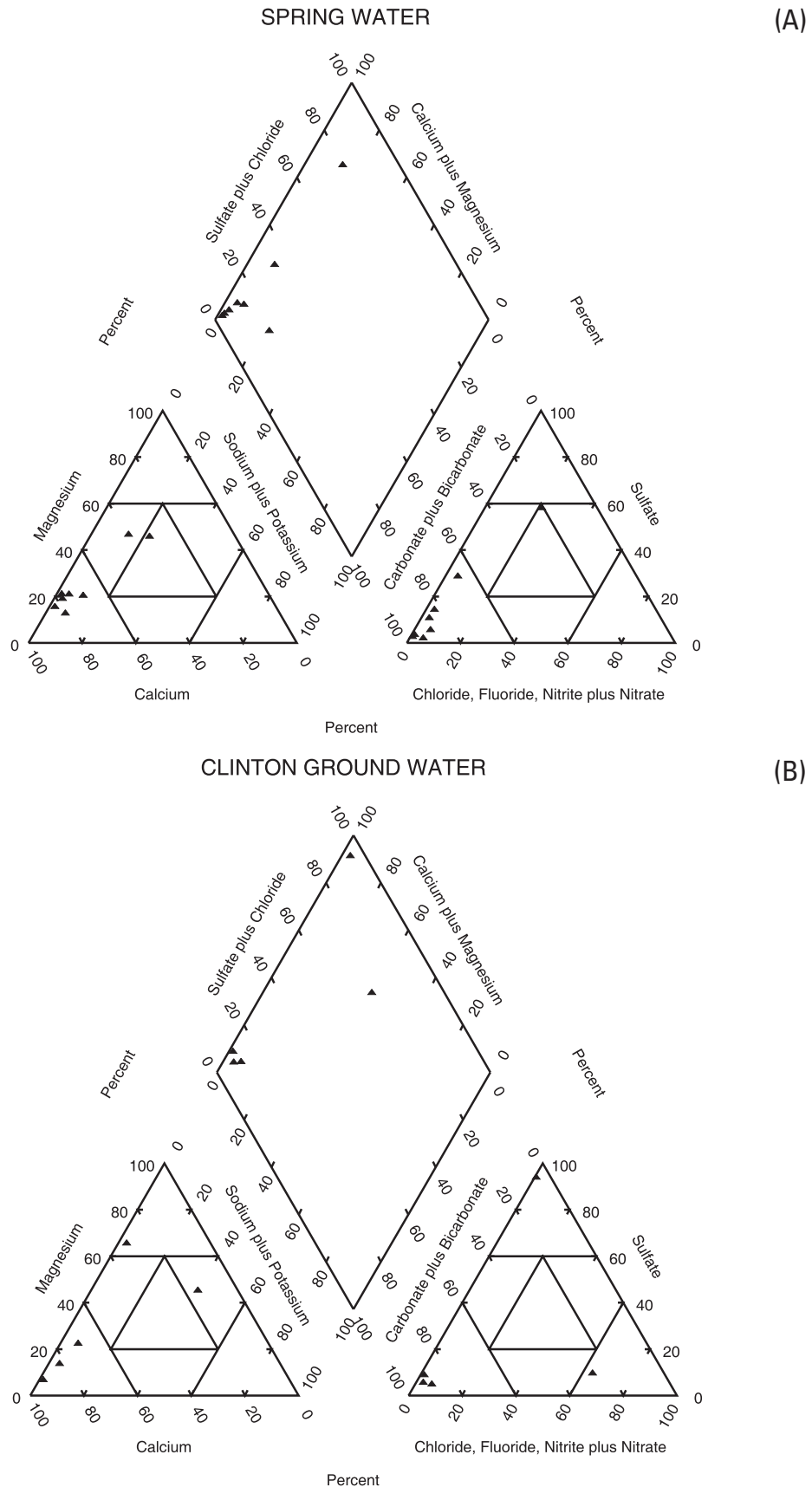


Figure 16. Relative geochemical composition of ground water from springs and wells, by hydrogeologic unit, Morgan County, West Virginia.

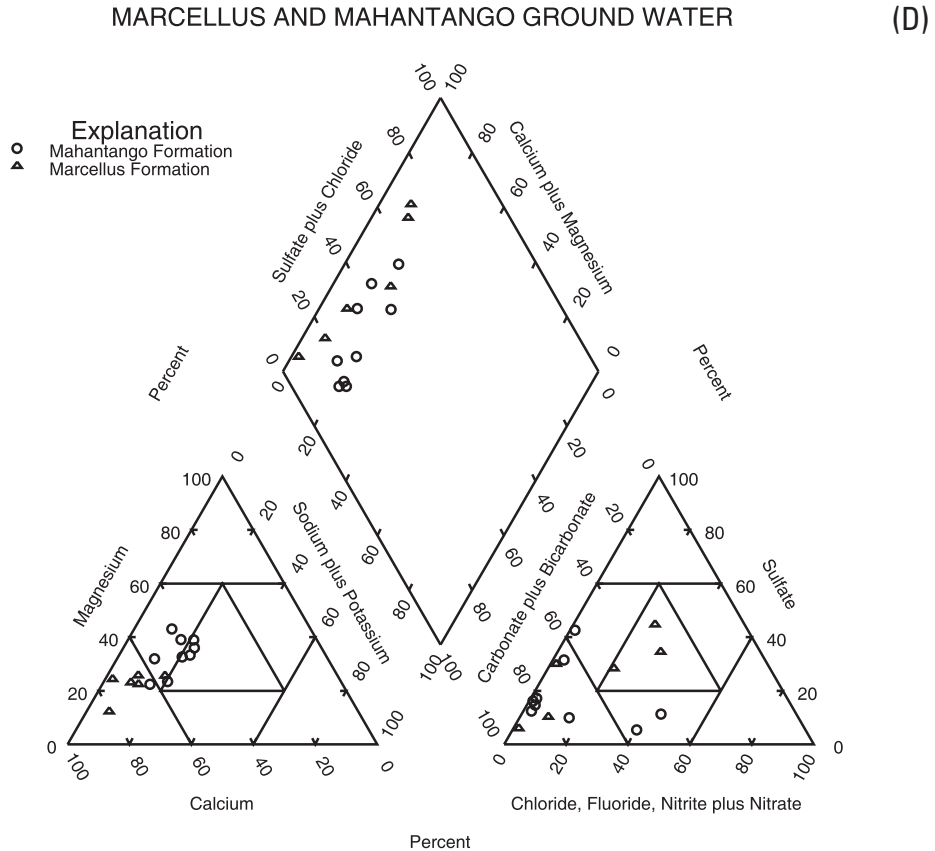
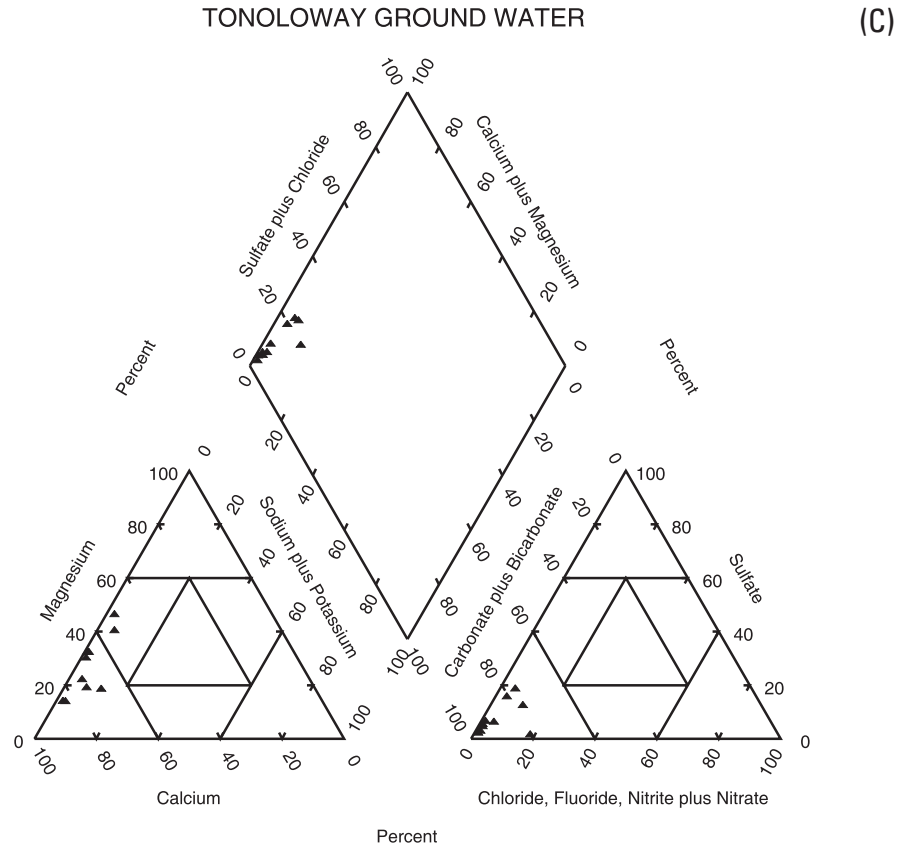


Figure 16. Relative geochemical composition of ground water from springs and wells, by hydrogeologic unit, Morgan County, West Virginia.—Continued

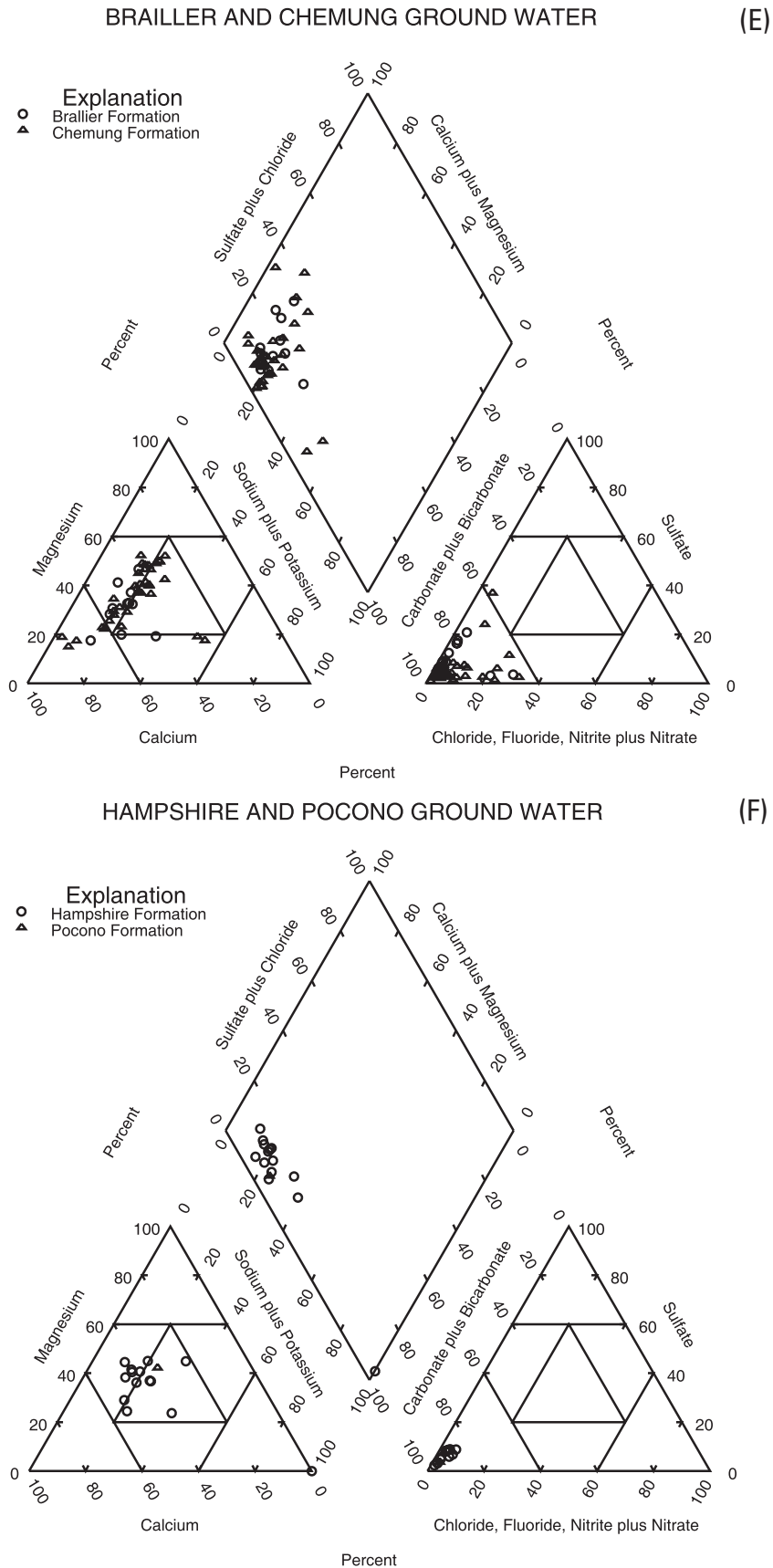


Figure 16. Relative geochemical composition of ground water from springs and wells, by hydrogeologic unit, Morgan County, West Virginia.—Continued

Iron and Manganese

Trace concentrations of iron and manganese occur naturally in many rocks and soils as various mineral assemblages. Martens (1945) reported deposits of limonite ($Fe_2O_3 \cdot H_2O$), an iron ore mineral occurring 6 mi west of Hedgesville, near the eastern boundary of Morgan County. The deposits are neither large nor of sufficiently high quality to be of contemporary commercial interest. Their occurrence does provide insight as to one potential source of large concentrations of iron in ground water. Concentrations of dissolved iron greater than the SMCL of 300 $\mu g/L$ typically precipitate when exposed to air or to DO in water and may cause turbidity, staining, and objectionable taste and color, and may corrode pipes, pumps, or other equipment (U.S. Environmental Protection Agency, 2005d). With the exception of the Tonoloway Formation and the one sample from the Pocono Group, ground water with concentrations of dissolved iron greater than 300 $\mu g/L$ was found in all formations throughout the study area (fig. 17).

Ground water with a concentration of dissolved iron of at least 300 $\mu g/L$ occurred in approximately 32 percent of the wells sampled (Appendix 2). The maximum concentration of dissolved iron in all ground-water samples was 24,100 $\mu g/L$ and the median was 42 $\mu g/L$. The maximum concentration of

dissolved iron in spring-water samples (site 93) was 2,010 $\mu g/L$. The concentration of dissolved iron in all other spring-water samples was less than the detection limit of 6 $\mu g/L$.

Dissolved manganese in excess of the SMCL of 50 $\mu g/L$ may result in black to brown color, black staining, and a bitter taste in the water (U.S. Environmental Protection Agency, 2005d). The maximum concentration of dissolved manganese in all ground-water samples was 3,920 $\mu g/L$ (fig. 18) and the median was 106.5 $\mu g/L$. The maximum concentration of dissolved manganese in spring-water samples was 212 $\mu g/L$. The median concentration of dissolved manganese from all spring-water samples was 0.60 $\mu g/L$. Ground water with concentrations of dissolved manganese greater than 50 $\mu g/L$ was found in all formations (fig. 18). Approximately 60 percent of the ground-water samples had concentrations of dissolved manganese greater than the SMCL (U.S. Environmental Protection Agency, 2005d).

The concentrations of iron and manganese are positively related (fig. 19). If the concentration of iron is small at a site, the concentration of manganese tends to be small. If the concentration of iron is large at a site, the concentration of manganese generally is large also.

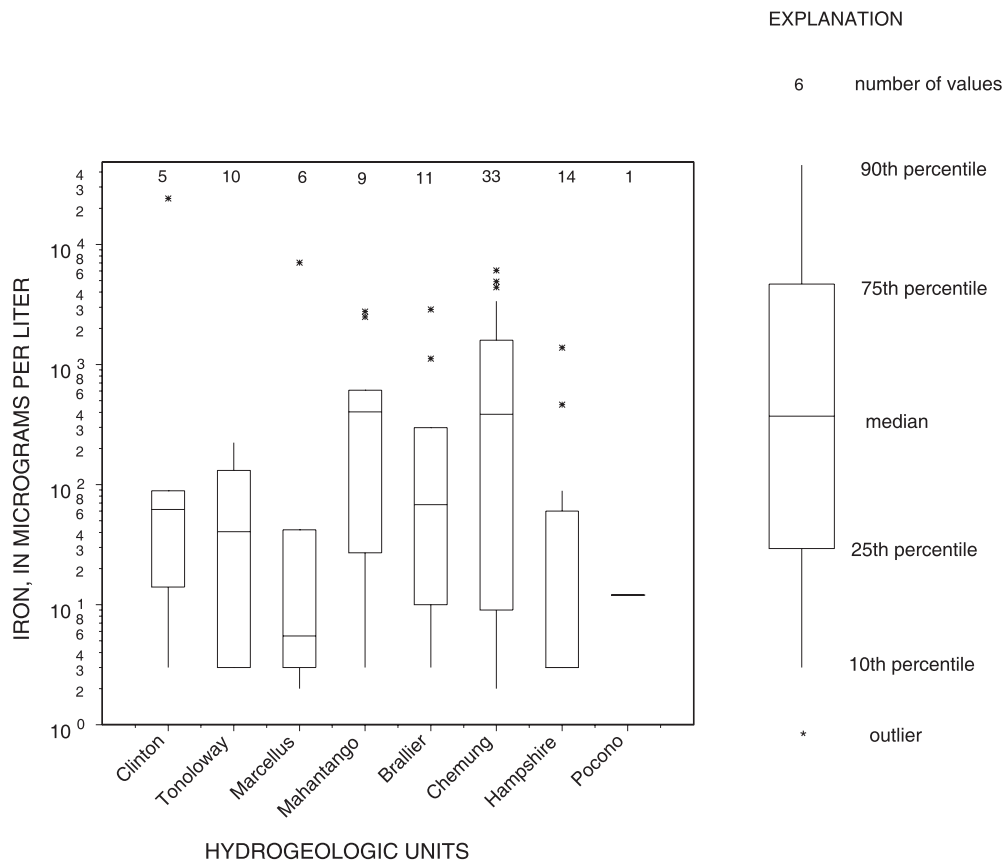


Figure 17. Distribution of concentrations of iron in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.

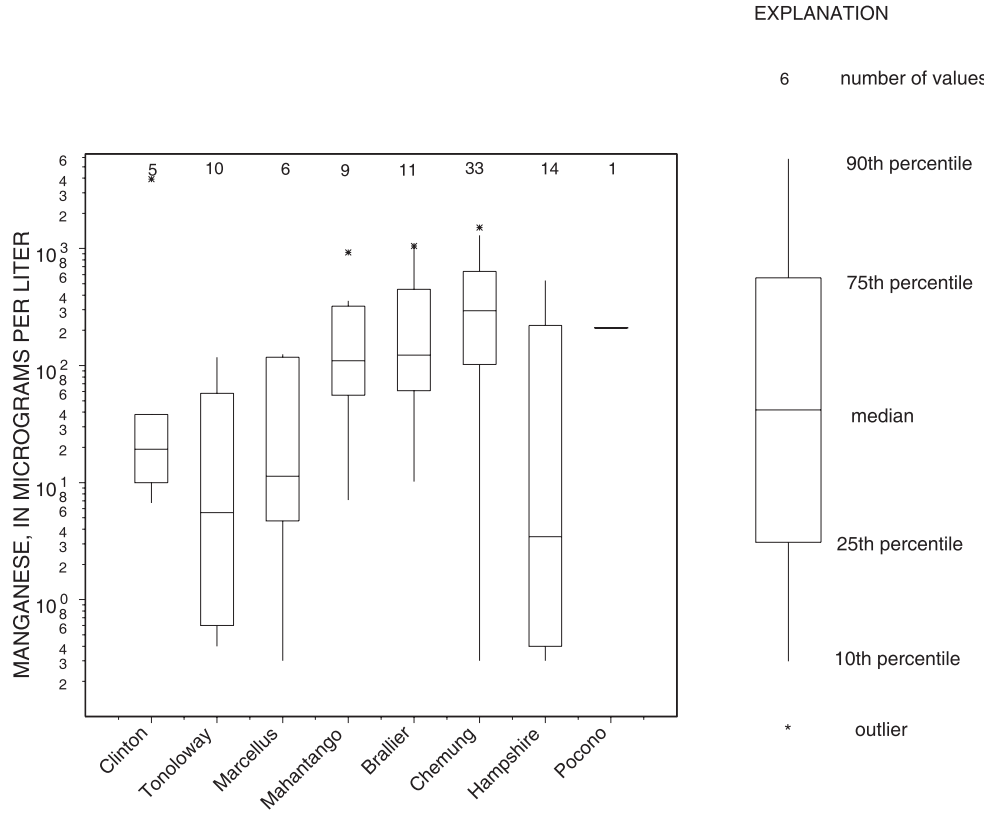
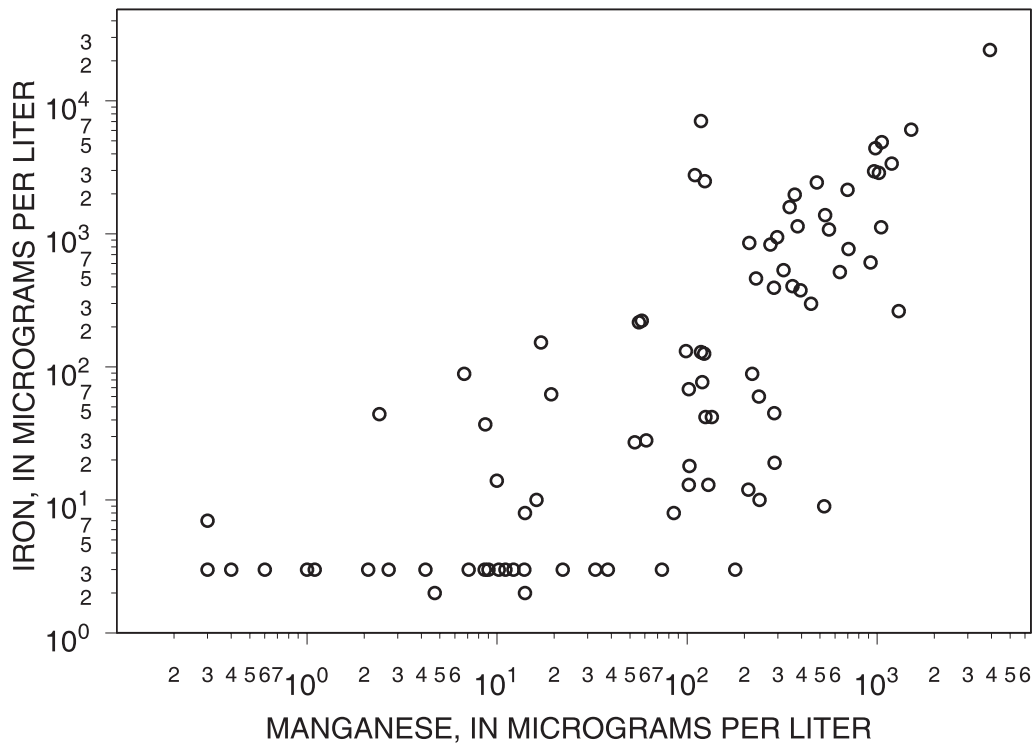


Figure 18. Distribution of concentrations of manganese in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.



Trace Elements

Trace elements occur in aquifer systems in which naturally occurring geochemical conditions are favorable for dissolving the elements from assemblages in the rock matrix. Trace elements sampled at a subset of sites were aluminum, arsenic, antimony, barium, cadmium, lead, nickel, strontium, and zinc.

The SMCL for aluminum (U.S. Environmental Protection Agency, 2005d) is a range from 50 to 200 $\mu\text{g/L}$. Aluminum may cause colored water, scaling, and sedimentation (U.S. Environmental Protection Agency, 2005d). The maximum concentration of aluminum sampled was 2,340 $\mu\text{g/L}$, with nine of the samples equal to or greater than the 50- $\mu\text{g/L}$ SMCL. Those samples were from the Brallier Formation, Chemung Group, and Clinton and Hampshire Formations. Five samples had concentrations less than the detection limit of 2 $\mu\text{g/L}$. Those samples occurred in the same formations and in the Tonoloway Formation. The median concentration of aluminum was 10 $\mu\text{g/L}$.

The MCL for antimony is 6 $\mu\text{g/L}$ (U.S. Environmental Protection Agency, 2005b). The potential health effects from exposure above that MCL are an increase in blood cholesterol and a potential for blood sugar to be decreased. Antimony was not detected in any sample at the minimum detection level of 0.20 $\mu\text{g/L}$.

The USEPA MCL for arsenic is 10 $\mu\text{g/L}$ (U.S. Environmental Protection Agency, 2005b). Concentrations of arsenic greater than the MCL may cause skin damage, problems with circulatory systems, or increased risk of cancer (U.S. Environmental Protection Agency, 2005b). The maximum concentration of arsenic sampled was 8 $\mu\text{g/L}$, with no sample concentrations greater than the 10 $\mu\text{g/L}$ MCL. Eighteen samples had concentrations less than the detection limit of 2 $\mu\text{g/L}$.

The MCL for barium is 200 $\mu\text{g/L}$ (U.S. Environmental Protection Agency, 2005b). Barium in concentrations exceeding the MCL may cause an increase in blood pressure (U.S. Environmental Protection Agency, 2005b). The maximum concentration of barium sampled was 390 $\mu\text{g/L}$, with nine sample concentrations greater than the 200- $\mu\text{g/L}$ MCL. Those samples came from the Brallier Formation, Chemung Group, and Clinton, Hampshire, and Tonoloway Formations. The smallest concentration of barium detected was 16.9 $\mu\text{g/L}$ and the median concentration was 76 $\mu\text{g/L}$.

The MCL for cadmium is 5 $\mu\text{g/L}$ (U.S. Environmental Protection Agency, 2005b). Concentrations of cadmium exceeding the MCL may cause kidney damage. The maximum concentration of cadmium sampled was 0.5 $\mu\text{g/L}$. None of the sample concentrations were greater than the 5- $\mu\text{g/L}$ MCL. The smallest concentration of cadmium was 0.02 $\mu\text{g/L}$ and the median concentration was 0.04 $\mu\text{g/L}$.

The MCL for lead is an action level of 15 $\mu\text{g/L}$ (U.S. Environmental Protection Agency, 2005b). For infants and children, lead may delay physical and mental development and

cause slight deficits in attention span and learning abilities. For adults, concentrations of lead exceeding the MCL may cause kidney problems or high blood pressure. The maximum concentration of lead detected was 1.7 $\mu\text{g/L}$ with 10 sample concentrations less than the detection limit of 0.05 $\mu\text{g/L}$. The median concentration of lead was 0.04 $\mu\text{g/L}$.

No MCL or SMCL currently (2006) exists for nickel (U.S. Environmental Protection Agency, 2005b and d). The maximum concentration of nickel detected was 30.6 $\mu\text{g/L}$, and the median concentration was 1.78 $\mu\text{g/L}$. Two samples had concentrations less than the detection limit. The three samples with the largest concentrations occurred in the Hampshire Formation.

No MCL or SMCL currently (2006) exists for elemental strontium (U.S. Environmental Protection Agency, 2005b and d). The maximum concentration of strontium detected was 5,520 $\mu\text{g/L}$. The smallest concentration of strontium was 8.87 $\mu\text{g/L}$ and the median concentration was 87.6 $\mu\text{g/L}$. The three samples with the largest concentrations were from the Hampshire Formation.

The SMCL for zinc is 500 $\mu\text{g/L}$ (U.S. Environmental Protection Agency, 2005d). Concentrations of zinc exceeding the SMCL may give a metallic taste to the water. The maximum concentration of zinc sampled was 696 $\mu\text{g/L}$, the only concentration greater than the 500- $\mu\text{g/L}$ SMCL; the sample was from the Marcellus Formation. The median concentration was 46.0 $\mu\text{g/L}$. Eleven samples had concentrations below the detection limit of 6-micrograms per liter.

Zinc and aluminum occur contemporaneously in the study area (fig. 20). The concentration of zinc is generally small at sites with a small concentration of aluminum. With little exception, the concentration of aluminum is large at site with a large concentration of zinc. A less strong positive relation exists between concentrations of strontium and lead.

Nutrients

All samples were analyzed for the following nutrients: ammonia, nitrate-plus-nitrite, nitrate, total phosphorus, and orthophosphate. The occurrence of nutrients in ground water at elevated concentrations frequently is caused by contamination from septic tank effluent, confined animal-feedlot operations, or fertilizer application. At elevated concentrations, certain nutrients may cause adverse effects on human health.

Nitrate and Nitrite

Small concentrations of nitrogen species occur naturally in rocks. Soils and biological matter contain larger concentrations of nitrogen species. In ground water, nitrate is usually the only nitrogen form of significance, although nitrogen species may occur at all oxidation states from -3 to +5, depending upon conditions (Hem, 1985).

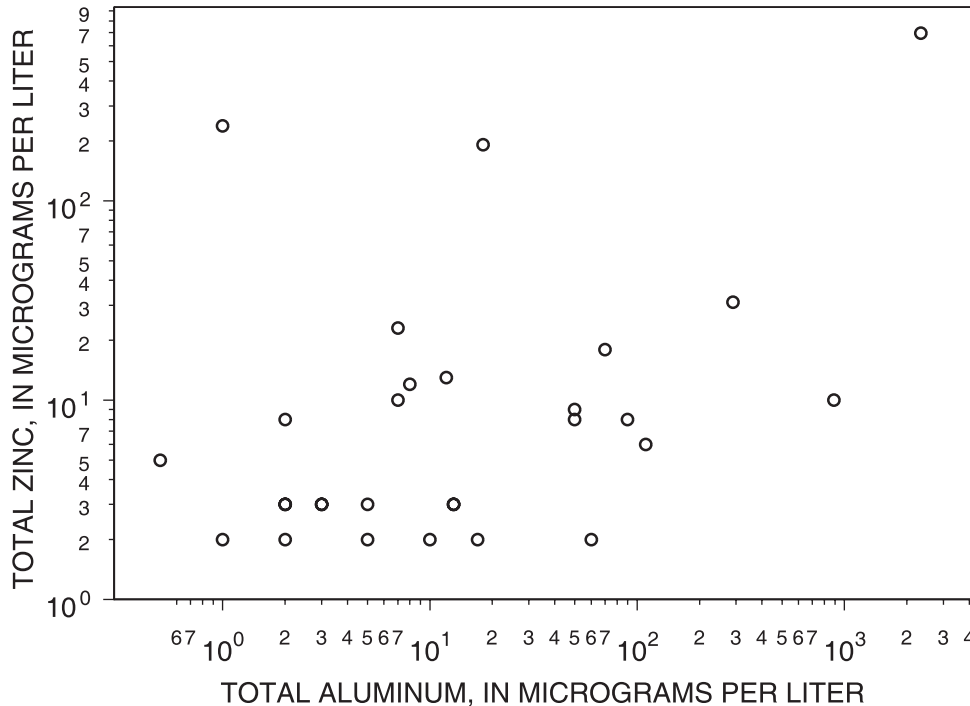


Figure 20. Relation between concentrations of zinc and aluminum in ground-water samples from Morgan County, West Virginia.

USEPA has established MCLs for nitrate of 10 mg/L as N and for nitrite of 1.0 mg/L as N (U.S. Environmental Protection Agency, 2005a). Infants below the age of 6 months may suffer from an acute condition known as methemoglobinemia (blue-baby syndrome) from water containing concentrations that exceed the MCLs. Nitrates and nitrites also have the potential to cause diuresis, increased starchy deposits, and hemorrhaging of the spleen from long-term exposures at concentrations exceeding the MCLs (U.S. Environmental Protection Agency, 2005b).

The largest concentration of nitrate-plus-nitrite as N was 2.6 mg/L, and 69 percent of the samples had concentrations less than the detection limit of 0.06 mg/L. The maximum concentration of nitrite as N, 0.042 mg/L, was the only value larger than the minimum reporting level (MRL), 0.008 mg/L, although there are several estimated values below the MRL.

Ammonia Species

The sources of ammonia species in natural water are assumed to be similar to those of other nitrogen species and typically are transformed to nitrate; however, ammonia is present in some natural waters (Hem, 1985). The largest concentration of ammonia as N detected during this study was

0.20 mg/L, and approximately 70 percent of the samples had concentrations less than the detection limit of 0.04 mg/L. No MCL or SMCL currently (2006) exists for ammonia in drinking water (U.S. Environmental Protection Agency, 2005b).

Phosphorous and Orthophosphate

Phosphorous occurs in a wide range of oxidation states in natural water; the most probable species are phosphate anions, complexes with metal ions, and colloidal particulate material (Hem, 1985). Phosphorous, as an element essential for animal metabolism, is always present in animal metabolic waste (Hem, 1985). As such, its presence in large concentrations may indicate contamination from animal or human waste. No MCL or SMCL currently (2006) exists for phosphorous or orthophosphate in drinking water (U.S. Environmental Protection Agency, 2005b).

The largest concentration of phosphorous detected during this study was 0.25 mg/L and more than 85 percent of the samples had concentrations less than the detection limit of 0.04 mg/L. The largest concentration of orthophosphate as P was 0.17 mg/L and more than 84 percent of the samples had concentrations less than the detection limit of 0.02 mg/L (U.S. Environmental Protection Agency, 2005b).

Radon-222

Radon-222 is a radioactive gas that occurs naturally as a decay product of uranium. Radon-222 gas is dissolved in ground water and degasses when the water is exposed to the open atmosphere. It can enter homes when water valves are opened, particularly during the use of showers. Rocks that are known to have a high uranium content include granitic rocks, some volcanic rocks, dark shales, sedimentary rocks that contain phosphate, and metamorphic rocks derived from these rocks (Otton and others, 1993). The accumulation and transport of radon in the subsurface depends upon the uranium content, grain size, permeability, and the extent of fracturing in the host rock. The rocks of Morgan County include dark shales and other sedimentary rocks that can contain uranium.

Radon-222 can destroy lung tissue and cause lung cancer (Otton and others, 1993) or can cause stomach cancer (U.S. Environmental Protection Agency, 2005c). In 1999, USEPA reinstated a proposed enforceable maximum contaminant level (PMCL) for radon in drinking water of 300 pCi/L, but with an alternate maximum contaminant level (AMCL) proposed at a concentration of 4,000 pCi/L. The drinking-

water standard that would apply for a community water system (CWS) would depend upon whether or not the State or CWS develops a radon Multimedia Mitigation (MMM) program. If a CWS serves 10,000 persons and either the State or the system has an approved MMM program, then the 4,000 pCi/L AMCL would apply; otherwise, the 300 pCi/L standard would apply. USEPA further proposes a nonenforceable Maximum Contaminant Level Goal (MCLG) of zero for radon-222 in drinking water (U.S. Environmental Protection Agency, 2005c). MCLGs are defined by USEPA as the level of a contaminant in drinking water below which there is no known or expected risk to health. MCLGs allow for a margin of safety and are nonenforceable public-health goals.

Radon-222 was detected in all wells sampled, ranging from 10 to 2,170 pCi/L, with a median value of 365 pCi/L. The two samples from the Clinton Formation had radon-222 concentrations much larger than the other formations, 1,330 and 2,170 pCi/L. Concentrations in 56 percent of the samples were greater than the PMCL of 300 pCi/L. Concentrations in none of the samples were greater than the proposed AMCL of 4,000 pCi/L (fig. 21).

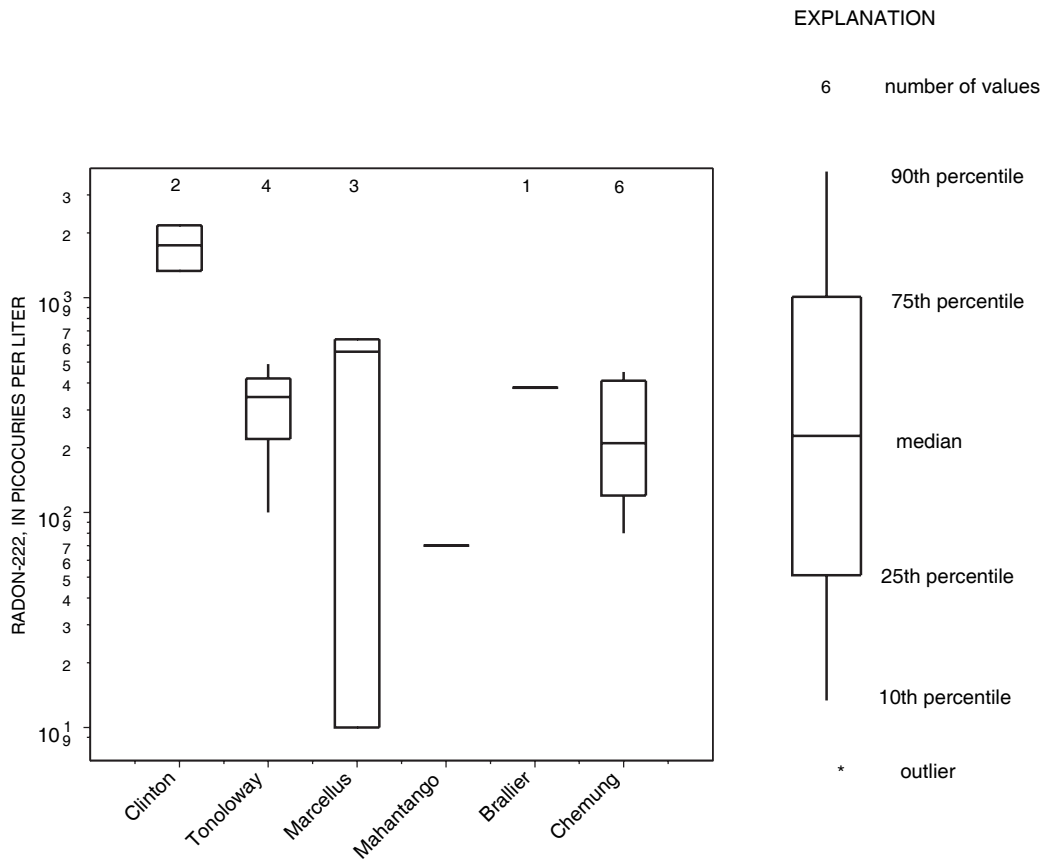


Figure 21. Distribution of concentration of radon-222 in ground-water samples from hydrogeologic units located in Morgan County, West Virginia.

Bacteria

Well and spring samples were tested for fecal coliform bacteria, total coliform bacteria, and *Escherichia coli* (*E. coli*). These microorganisms occur in the feces of warmblooded animals, including humans. The presence of these microorganisms indicates the presence of fecal contamination and the possible presence of other harmful bacteria, viruses, or other pathogens into the ground-water system. Certain strains of *E. coli* are of particular concern because they can cause health problems such as fever, diarrhea, and, in certain instances, kidney failure (U.S. Centers for Disease Control and Prevention, 2005). USEPA has established an MCLG for total coliforms (including fecal coliform and *E. coli*) of zero colonies per 100 milliliters (col/100 mL) for public drinking water supply systems (U.S. Environmental Protection Agency, 2005a).

Ninety-one wells in the study area were tested for coliform bacteria. Fecal coliform bacteria were detected in 5 ground-water samples (5 percent), total coliform bacteria were detected in 42 ground-water samples (46 percent), and *E. coli* were detected in 2 ground-water samples (2 percent).

Approximately 69 percent of the wells sampled were constructed after 1983 and constructed with a grout surface seal (West Virginia Bureau for Public Health, 1984). West Virginia Bureau for Public Health (1984) requires installation of a concrete pad around the casing at the surface but does not require grouting the entire length of the casing. Proper surface seals and grouting of wells helps to reduce the potential for transport of fecal contaminants along the casing from the surface to the aquifer. The casings of most of the wells sampled appeared to be in good condition at the surface, although no down-hole observations were possible. Additionally, most of the wells were located at a distance from potential fecal-material sources such as barnyard manure or pet waste. Proper siting of wells with respect to septic systems can reduce or eliminate the potential of microbial transport through the aquifer from that source.

The occurrence of total coliform bacteria was compared with geologic unit, age of well, depth of well, the ratio of well depth to casing length, and aquifer transmissivity as determined during the recovery phase of the aquifer test. No significant correlation was found among the occurrence of total coliform bacteria and any of these characteristics. Well-site proximities to septic systems were not evaluated as part of this study.

Mathes (2000) studied the relation of bacteria to septic systems in Berkeley County, W. Va., and found no apparent relation among septic-system density and the concentration of bacterial colonies detected. This lack of correlation

was attributed to the structural complexity of the limestone aquifers in the study area. The geology in Morgan County is equally complex. Soils are generally shallow, and all of the bedrock units have undergone faulting, folding, and fracturing. Septic systems sited on shallow soils may result in percolation through underlying fractures, where contamination may be transported rapidly into the aquifer. Mathes (2000) found little difference in the frequency of total coliform bacteria between shallow wells and deep wells. Similarly, no correlation was found during this study. Mathes (2000) also found that the greatest concentrations of *E. coli* and fecal coliform bacteria were in shallow wells. No such correlation was found during this study; however, evaluation of well-completion reports for Morgan County indicates that well production in the domestic wells frequently consists of two or more fracture zones. The uppermost fracture zone is frequently at a depth of less than 100 ft below land surface, and produces 1 to 2 gal/min, inadequate for a domestic supply well. Drilling is continued until another fracture zone is penetrated, providing a combined production adequate to supply domestic needs. The upper zones are rarely grouted; therefore, the potential exists for contamination from shallower zones even in deep wells.

The ratio of well depth to casing length was developed and correlated with the occurrence of total coliform bacteria. No significant correlation exists; however, the sites with the largest total coliform bacteria concentration had a ratio of total well depth to casing length less than or equal to 0.35 (fig. 22). The most contaminated site with a ratio exceeding 0.35 was a newly constructed well, which may have had residual contamination from drilling. In highly folded, faulted, and fractured terrain, longer grouted casings may be needed to protect underlying aquifers.

Eight springs were sampled for the presence of bacteria. No bacteria were detected in two of the eight springs. Total coliform bacteria were detected in six of the springs, and *E. coli* were detected at two spring sites.

Natural springs are open to the surrounding environment and are therefore susceptible to microbial contamination by feces from wildlife or domestic animals that frequent the area. Properly constructed spring houses can protect a water source from fecal contamination. The two springs that were free from coliform bacteria have enclosed spring structures. With the exception of one free-flowing, natural spring, all of the springs tested were protected by spring houses or spring boxes. The structures varied significantly from site to site. At one site, frogs were observed in a spring house that had an unscreened discharge pipe that could serve as an opening for other animals to enter. No inventory of potential sources of contamination was done during the course of this study.

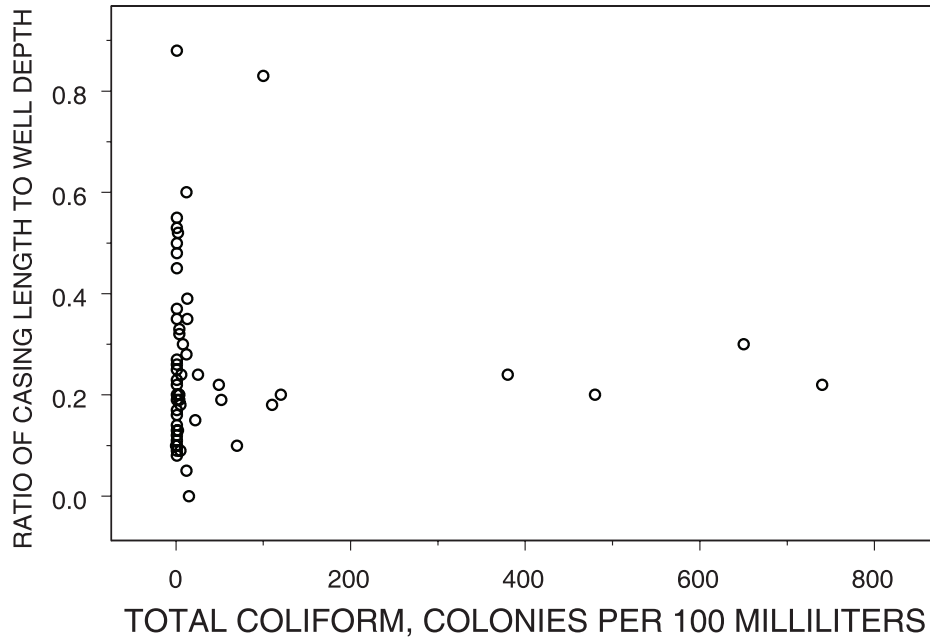


Figure 22. Relation between the ratio of casing length to well depth and the occurrence of bacteria in wells located in Morgan County, West Virginia.

Summary and Conclusions

With few exceptions, municipal and private users in the study area depend upon ground water for their water supplies. Aquifers in Morgan County are composed of carbonate, sandstone, and shale rocks of ages ranging from Late Ordovician to Early Mississippian. The rocks are characterized by thin to thick bedded formations. Interbedding occurs among the various limestones, shales, sandstones, and siltstones that includes calcareous shale and sandstone and cherty limestone and shale. The hydrogeology is complicated by the structural complexity typical of the Valley and Ridge Physiographic Province. That structural complexity consists of a series of steeply dipping north-south trending anticlines, synclines, and thrust faults that are intersected by cross-faulting in several locations throughout the study area.

Static ground-water altitudes decrease from recharge areas in the mountains to the valley floors. Along the valley floors, altitudes decrease from the southern part of the county to the Potomac River in the north. Thus, ground water ultimately flows in a northerly direction. Small drainages have developed where cross-faulting or jointing occurs and these drainages provide local recharge zones. Static water levels may vary among nearby or adjacent sites by as much as 200 ft in some localities because of the complexity of the geologic stratigraphy and structure. This variability also may indicate the presence of separate or otherwise isolated aquifers at different depths because of the local structural complexity.

Transmissivities were determined for aquifers throughout the study area. Transmissivities ranged from 2 to 1,420 ft²/d. The largest transmissivities generally occurred at or near contacts between geologic units (at bedding planes) or in the

presence of thrust-faults or cross-faults. Small transmissivities (less than or equal to 5 ft²/d) occurred in all aquifer units except the Marcellus Formation, which had transmissivities ranging from 54 to 327 ft²/d and had the largest median transmissivity at 128 ft²/d. The largest transmissivity recorded was 1,490 ft²/d from the Chemung Group, but near a bedding contact with the Brallier Formation. Production from wells within the study area is typically from one or two fracture zones.

Water quality and geochemistry are variable throughout the county and within aquifer units. Geochemically, the ground water ranges from nearly 100 percent sodium-bicarbonate water to nearly 100 percent calcium-magnesium-sulfate-chloride water as end-members. Most of the water tested is calcium-bicarbonate water, which is typical of water from a limestone aquifer. Geochemical mixing has occurred among or within aquifer units.

Water-quality characteristics for which values were greater than U.S. Environmental Protection Agency Maximum Contaminant Levels or Secondary Maximum Contaminant Levels include radon-222, pH, turbidity, iron, manganese, aluminum, and total- and fecal-coliform and *E. coli* bacteria. Concentrations of radon-222 were greater than the 300 pCi/L Proposed Maximum Contaminant Level in more than half of the wells tested. Total coliform occurred in 42 wells, fecal coliform in 5 wells, and *E. coli* in 2 wells. More than 35 percent of the wells tested had turbidity greater than 5 NTU. The largest measurement was 330 NTU. Pumping stress on aquifers with low production tended to increase turbidity. Specific conductance (SC) ranged from 44 to 737 μ S/cm. The median SC was 240 μ S/cm. Hardness as CaCO₃ ranged from 16 to 400 mg/L. The median hardness was 98 mg/L.

Acknowledgments

The authors thank Danny Evans of the West Virginia Conservation Agency and the Eastern Panhandle Conservation District for their cooperation in this study. We also gratefully acknowledge the Morgan County Rural Water Committee and the Morgan County Planning Commission for their assistance and West Virginia University for their collaboration during the course of the study. We particularly thank the many individuals, establishments, and authorities in the county that permitted access to their wells for collection of hydraulic and water-quality data.

References Cited

- American National Standards Institute, 2004, Standard test method for radon in drinking water: Annual Book of ASTM Standards, sec. 11, v. 11.02, p. 662-664.
- American Public Health Association, 1998, Standard methods for the examination of water and wastewater (20th ed.): Washington, D.C., American Public Health Association, American Water Works Association, and Water Environment Federation, p. 3-37 – 3-43.
- Brownlow, A.H., 1979, *Geochemistry*: Englewood Cliffs, N.J., Prentice-Hall, p. 289-298.
- Cardwell, D.H., Ervin, R.B., and Woodward, H.P., 1968, *Geologic map of West Virginia*: West Virginia Geological and Economic Survey, 2 sheets, scale 1:250,000.
- Clark, W.E., Chisholm, J.L., and Frye, P.M., 1976, *Water resources of the upper New River basin, West Virginia*: West Virginia Geological and Economic Survey and West Virginia Department of Natural Resources, Division of Water Resources, River Basin Bulletin 4.
- Cooper, C.K., and Hanna, E.A., 1988, *Field water-quality data collected at gaging stations in West Virginia, 1945-1975*: U.S. Geological Survey Open-File Report 86-420K, 426 p.
- Doll, W.L., Meyer, G., and Archer, R.J., 1963, *Water resources of West Virginia*: West Virginia Department of Natural Resources, Division of Water Resources report, 134 p.
- Driscoll, F.G., 1986, *Groundwater and wells*: St. Paul, Minn., Johnson Filtration Systems Inc., 1,089 p.
- Durfor, C.N., and Becker, Edith, 1964, *Public water supplies of the 100 largest cities in the United States, 1962*: U.S. Geological Survey Water-Supply Paper 1812, 364 p.
- Edmunds, W.M., Bath, A.H., and Miles, D.L., 1982, *Hydrochemical evolution of the East Midlands Triassic sandstone aquifer, England*: *Geochimica et Cosmochimica Acta*, v. 46, p. 2069-2081.
- Erskine, H.M., 1948, *Principal springs of West Virginia, a report on the location, discharge, and temperature of the principal springs of West Virginia*: West Virginia Conservation Commission and West Virginia State Water Commission, 51 p.
- Faires, L.M., 1993, *Methods of analysis by the U.S. Geological Survey National Water Quality Laboratory—Determination of metals in water by inductively coupled plasma-mass spectrometry*: U.S. Geological Survey Open-File Report 92-634, 28 p.
- Fenneman, N.M., 1946, *Physical Divisions of the United States*: U.S. Geological Survey, scale 1:7,000,000.
- Ferrell, G.M., 1987, *West Virginia ground-water quality*: U.S. Geological Survey Open-File Report 87-0761, 11 p.
- Fishman, M.J., ed., 1993, *Methods of analysis by the U.S. Geological Survey National Water Quality Laboratory—Determination of inorganic and organic constituents in water and fluvial sediments*: U.S. Geological Survey Open-File Report 93-125, 217 p.
- Fishman, M.J., and Friedman, L.C., 1989, *Methods for determination of inorganic substances in water and fluvial sediments*: U.S. Geological Survey Techniques of Water-Resources Investigations, book 5, chap. A1, 545 p.
- Freeze, R.A., and Cherry, J.A., 1979, *Groundwater*: Englewood Cliffs, N.J., Prentice-Hall, 604 p.
- Garbarino, J.R., 1999, *Methods of analysis by the U.S. Geological Survey National Water Quality Laboratory—Determination of dissolved arsenic, boron, lithium, selenium, strontium, thallium, and vanadium using inductively coupled plasma-mass spectrometry*: U.S. Geological Survey Open-File Report 99-093, 31 p.
- Garbarino, J.R., and Struzeski, T.M., 1998, *Methods of analysis by the U.S. Geological Survey National Water Quality Laboratory—Determination of elements in whole-water digests using inductively coupled plasma-optical emission spectrometry and inductively coupled plasma-mass spectrometry*: U.S. Geological Survey Open-File Report 98-165, 101 p.
- Garrels, R.M., and Christ, C.L., 1965, *Solutions, minerals, and equilibria*: San Francisco, Calif., Freeman, Cooper & Company, p. 379-402.
- Grimsley, G.P., 1916, *Jefferson, Berkeley, and Morgan County report*: West Virginia Geological Survey County Report, 644 p., 3 maps.

- Heath, R.C., 1983, Basic ground-water hydrology: U.S. Geological Survey Water-Supply Paper 2220, 85 p.
- Hem, J.D., 1985, Study and interpretation of the chemical characteristics of natural water: U.S. Geological Survey Water-Supply Paper 2254, 264 p.
- Hobba, W.A., Jr., 1985, Water in Hardy, Hampshire, and western Morgan Counties, West Virginia: West Virginia Geological and Economic Survey Environmental Geology EGB-19, 91 p.
- Hobba, W.A., Jr., Fisher, D.W., Pearson, F.J., Jr., and Chermers, J.C., 1979, Hydrology and geochemistry of thermal springs of the Appalachians: U.S. Geological Survey Professional Paper 1044-E, 36 p.
- Hobba, W.A., Jr., Friel, E.A., and Chisholm, J.L., 1972, Water resources of the Potomac River basin, West Virginia: West Virginia Geological and Economic Survey River Basin Bulletin 3, 110 p.
- Hoffman, J.F., 1962, Water and the future of West Virginia: West Virginia Division of Water Resources and West Virginia Institute of Technology Bulletin 1, 16 p.
- Holmes, D.E., ed., 1995, Counties in West Virginia Blue Book—1995: Chapman Printing, vol. 77, p. 731-846.
- Hsieh, P.A., and Shapiro, A.M., 1996, Hydraulic characteristics of fractured bedrock underlying the FSE well field at the Mirror Lake site, Grafton County, New Hampshire, *in* Morganwalp, D.W., and Aronson, D.A., eds., U.S. Geological Survey Toxic Substances Hydrology Program—Proceedings of the technical meeting, Colorado Springs, Colo., September 20-24, 1993: U.S. Geological Survey Water-Resources Investigations Report 94-4015.
- Ice, J.R., comp., 1933, New descriptive atlas of West Virginia: Clarksburg, W.Va., Clarksburg Publishing Company, 60 p.
- Kozar, M.D., and Mathes, M.V., 2001, Aquifer characteristics data for West Virginia: U.S. Geological Survey Water-Resources Investigations Report 01-4036, 74 p.
- Kulander, B.R., Lessing, P., Dean, S.L., and Kulander, C.S., 1995, Geology of Stotlers Crossroads quadrangle, Berkeley and Morgan Counties, West Virginia: West Virginia Geological and Economic Survey Open-File Report OF9504, 1:24,000.
- Lessing, P., Dean, S.L., and Kulander, B.R., 1997, Geology of Great Cacapon and Bellegrave quadrangles, Morgan County, West Virginia: West Virginia Geological and Economic Survey Open-File Report OF9701, 1:24,000.
- Lessing, P., Hobba, W.A., Jr., Dean, S.L., and Kulander, B.R., 1991, Relations between warm springs and geology delineated by side-looking airborne-radar imagery in eastern West Virginia: U.S. Geological Survey Water-Resources Investigations Report 88-4096, 44 p.
- Martens, J.H.C., 1945, Fifty common rocks and minerals of West Virginia: West Virginia Geological Survey Bulletin 9, 42 p.
- Mathes, M.V., 2000, Relation of bacteria in limestone aquifers to septic systems in Berkeley County, West Virginia: U.S. Geological Survey Water-Resources Investigations Report 00-4229, 12 p.
- Mathes, M.V., Jr., Kozar, M.D., and Brown, D.P., 1998, Summary of ground-water quality in West Virginia: West Virginia Division of Environmental Protection, Office of Water Resources, Ground-Water Program, 54 p.
- McColloch, J.S., 1986, Springs of West Virginia: West Virginia Geological and Economic Survey, v. V06A, 493 p.
- McColloch, J.S., and Lessing, Peter, 1980, Land use statistics for West Virginia: West Virginia Geological and Economic Survey Environmental Geology Bulletin 18A, 62 p.
- McCoy, K.J., Podwysocki, M.H., Crider, E.A., and Weary, D.J., 2005a, Fracture trace map and single-well aquifer test results in a carbonate aquifer in Berkeley County, West Virginia: U.S. Geological Survey Open-File Report 2005-1040, accessed October 26, 2005, at <http://pubs.usgs.gov/of/2005/1040/>
- McCoy, K.J., Podwysocki, M.H., Crider, E.A., and Weary, D.J., 2005b, Fracture trace map and single-well aquifer test results in a carbonate aquifer in Jefferson County, West Virginia: U.S. Geological Survey Open-File Report 2005-1407, accessed January 19, 2006, at <http://pubs.usgs.gov/of/2005/1407/>
- Myers, D.N., and Wilde, F.D., eds., 2003, Biological indicators (3d ed.): U.S. Geological Survey Techniques of Water-Resources Investigations, book 9, chap. A7, 163 p.
- National Oceanic and Atmospheric Administration, 2002, Monthly normals of temperatures, precipitation, and heating and cooling degree days 1971-2000: Asheville, N.C., U.S. Department of Commerce, National Climatic Data Center, Climatology of the United States no. 81, 46-West Virginia, 23 p.
- Otton, J.K., Gundersen, L.C.S., and Schumann, R.R., 1993, The geology of radon: U.S. Geological Survey General Interest Publication, 29 p.

- Patton, C.J., and Truitt, E.P., 1992, Methods of analysis by the U.S. Geological Survey National Water Quality Laboratory—Determination of total phosphorus by a Kjeldahl digestion method and an automated colorimetric finish that includes dialysis: U.S. Geological Survey Open-File Report 92-146, 39 p.
- Perry, L.D., Costain, J.K., and Geiser, P.A., 1979, Heat flow in western Virginia and a model for the origin of thermal springs in the Appalachians: *Journal of Geophysical Research*, v. 84, no. B12, p. 6875-6883.
- Reeves, Frank, 1932, Thermal springs of Virginia: Virginia Geological Survey, Bulletin 36, 56 p.
- Reger, D.B., 1924, County reports—Mineral and Grant Counties: West Virginia Geological and Economic Survey, 866 p.
- Schad, H., and Teutsch, G., 1994, Effects of the investigation scale on pumping test results in heterogeneous porous aquifers: *Journal of Hydrology*, v. 159, no. 1-4, p. 61-77.
- Schwietering, J.F., 1981, Brief description of ground water conditions and aquifers in West Virginia with annotated bibliography: West Virginia Geological and Economic Survey Open File Report OF-8102, 121 p.
- Shultz, R.A., Hobba, W.A., and Kozar, M.D., 1995, Geohydrology and ground-water quality of Berkeley County, West Virginia, with emphasis on the carbonate area: U.S. Geological Survey Water-Resources Investigations Report 93-4073, 88 p.
- Siddiqui, S.H., and Parizek, R.R., 1972, Application of non-parametric statistical tests in hydrogeology: *Ground Water*, v. 10, no. 2, p. 26-31.
- Teideman, C.R., and Hsieh, P.A., 2001, Assessing open-well aquifer test in fractured, crystalline rock: *Ground Water*, v. 39, no. 1, p. 68-78.
- Theis, C.V., 1935, The lowering of the piezometer surface and the rate and discharge of a well using ground-water storage: *American Geophysical Union Transactions*, v. 16, p. 519-524.
- Toth, J., 1963, A theoretical analysis of groundwater flow in small drainage basins: *Journal of Geophysical Research*, v. 68, p. 4795-4812.
- U.S. Centers for Disease Control and Prevention, 2005, *Escherichia coli* O157:H7, accessed October 24, 2005, at http://www.cdc.gov/ncidod/dbmd/diseaseinfo/escherichiacoli_g.htm#What%20illness%20does%20E.%20coli%20157:H7%20cause
- U.S. Department of Commerce, 1991, Comparison of 1980 and 1990 population data West Virginia and Counties: U.S. Census Bureau, Washington, D.C., 9 p.
- U.S. Department of Commerce, 2005, State and County quick-facts: U.S. Census Bureau, accessed October 26, 2005, at <http://quickfacts.census.gov/qfd/states/54/54065.html>
- U.S. Environmental Protection Agency, 2005a, Consumer factsheet on nitrates/nitrites, accessed October 24, 2005, at http://www.epa.gov/safewater/contaminants/dw_contamfs/nitrates.html
- U.S. Environmental Protection Agency, 2005b, List of drinking water contaminants and MCLs, accessed October 24, 2005, at <http://www.epa.gov/safewater/mcl.html>
- U.S. Environmental Protection Agency, 2005c, Proposed radon in drinking water rule, EPA 815-F-99-009, accessed October 24, 2005, at <http://www.epa.gov/safewater/radon/proposal.html>
- U.S. Environmental Protection Agency, 2005d, Secondary Drinking Water Regulations, guidance for nuisance chemicals, accessed October 24, 2005, at <http://www.epa.gov/safewater/consumer/2ndstandards.html>
- West Virginia Bureau for Public Health, 1984, Water well design standards, ch. 16-1, series 3: West Virginia Bureau for Public Health report SW-255, 9 p.
- Wilde, F.D., and Radtke, D.B., eds., 1998, Field measurements, in National field manual for the collection of water-quality data: U.S. Geological Survey Techniques of Water-Resources Investigations, book 9, chap. A6, [variously paged].
- Wilde, F.D., Radtke, D.B., Gibs, Jacob, and Iwatsubo, R.T., eds., 1999, Collection of water samples, in National field manual for the collection of water-quality data: U.S. Geological Survey Techniques of Water-Resources Investigations, book 9, chap. A4, [variously paged].
- Winograd, I.J., and Robertson, F.N., 1982, Deep oxygenated ground water—Anomaly or common occurrence?: *Science*, v. 216, p. 1227-1230.
- Wyrick, G.G., and Borchers, J.W., 1981, Hydrologic effects of stress-relief fracturing in an Appalachian valley: U.S. Geological Survey Water-Supply Paper 2177, 51 p.

Appendixes 1 through 3

Appendix 1. Well-construction and aquifer properties for selected wells, Morgan County, West Virginia.

[USGS, U.S. Geological Survey; ft bls, feet below land surface; ft, foot; in., inch; Hydrogeologic unit codes: 354CLNN, Clinton Group; 35ITNLY, Tomoloway Formation; 344MRCL, Marcellus Formation; 344MNNG, Mahantango Formation; 34IBRLR, Brallier Formation; 341CMNG, Chemung Group; 341HMPR, Hampshire Formation; 337POCN, Pocomo Group, gal/min, gallons per minute; (gal/min)/ft, gallons per minute per foot of drawdown; ft²/d, foot squared per day; --, not determined. See figure 9 for well locations]

Site number	USGS station name	Well depth (ft bls)	Length of casing (ft)	Casing diameter (in.)	Land-surface altitude (ft)	Static water level (ft bls)	Date water level measured	Hydrogeologic unit	Discharge (gal/min)	Specific capacity ((gal/min)/ft)	Transmissivity average (ft ² /d)
1	Mrg-0090	142	11	6	560	63.41	7-6-2004	344MNNG	--	0.10	6
2	Mrg-0178	225	59.4	6	801	103.02	5-10-2005	341CMNG	3.6	.10	3
3	Mrg-0146	120	40	6	745	34.61	11-3-2004	34IBRLR	5.7	.45	113
4	Mrg-0114	320	118	6	500	106.05	8-11-2004	341CMNG	--	.13	6.5
5	Mrg-0124	320	--	6	730	50.6	8-26-2004	341CMNG	--	3.75	317
6	Mrg-0104	140	39	6	780	48.85	8-2-2004	34IBRLR	--	1.00	96
7	Mrg-0122	145	--	6	680	78.9	8-25-2004	341CMNG	--	--	--
8	Mrg-0144	300	105	6	590	8.77	11-2-2004	34HMPR	3.2	.12	6
9	Mrg-0130	220	60	6	864	154.92	9-2-2004	354CLNN	7.8	1.05	206
10	Mrg-0095	140	--	6	710	26.59	7-12-2004	341CMNG	5.3	.91	29
11	Mrg-0139	160	83	6	744	35.02	10-19-2004	341CMNG	3.1	3.16	211.5
12	Mrg-0136	163	--	6	743	50.1	9-22-2004	341CMNG	4	--	--
13	Mrg-0142	160	42	6	797	43.17	10-21-2004	341CMNG	8.2	.90	92.5
14	Mrg-0145	180	100	6	751	23.78	10-19-2004	341CMNG	7.5	.33	21
15	Mrg-0135	177	40	6	870	90.75	9-22-2004	341CMNG	5.2	.49	55
16	Mrg-0156	160	39	6	589	45.43	4-13-2005	344MNNG	4.8	.53	32
17	Mrg-0094	200	20	6	480	24.93	7-12-2004	341CMNG	10.2	1.10	122
18	Mrg-0161	365	78.65	6	484	16.98	4-19-2005	35ITNLY	10.8	1.42	138
19	Mrg-0137	269	40	6	847	37.91	9-23-2004	341CMNG	1.5	.17	5
20	Mrg-0143	200	40	6	561	51.08	11-2-2004	344MNNG	4.8	.08	2
21	Mrg-0140	560	49	6	908	103.02	10-20-2004	341CMNG	5.7	.09	3
22	Mrg-0092	405	40	6	840	67.73	7-7-2004	341CMNG	9	.50	39.5
23	Mrg-0149	100	60	6	567	37.31	11-4-2004	344MNNG	2.6	1.43	131
24	Mrg-0147	225	40	6	532	62.9	11-3-2004	344MNNG	3.7	.36	108
25	Mrg-0112	120	59	6	680	20	8-10-2004	341CMNG	--	1.05	168
26	Mrg-0138	120	100	6	552	48.04	9-23-2004	344MNNG	6.1	4.61	577
27	Mrg-0141	150	30	6	586	34	10-20-2004	344MNNG	1.4	.12	6
28	Mrg-0176	265	59.0	6	731	79.71	5-6-2005	34HMPR	9	.86	74
29	Mrg-0159	350	--	6	768	--	4-15-2005	354CLNN	3	--	--
30	Mrg-0091	205	80	6	680	77.95	7-7-2004	341CMNG	5.2	.15	8
31	Mrg-0179	122	19.1	6	722	14.85	5-11-2005	344MRCL	5.6	.40	54

Appendix 1. Well-construction and aquifer properties for selected wells, Morgan County, West Virginia. — Continued

[USGS, U.S. Geological Survey; ft bls, feet below land surface; ft, foot; in., inch; Hydrogeologic unit codes: 354CLNN, Clinton Group; 351TNLY, Tomoloway Formation; 344MRCL, Marcellus Formation; 344MNNG, Mahantango Formation; 341BRLR, Brallier Formation; 341CMNG, Chemung Group; 341HMPR, Hampshire Formation; 337POCN, Pocomo Group, gal/min, gallons per minute; (gal/min)/ft, gallons per minute per foot of drawdown; ft²/d, foot squared per day; --, not determined. See figure 9 for well locations]

Site number	USGS station name	Well depth (ft bls)	Length of casing (ft)	Casing diameter (in.)	Land-surface altitude (ft)	Static water level (ft bls)	Date water level measured	Hydrogeologic unit	Discharge (gal/min)	Specific capacity ((gal/min)/ft)	Transmissivity average (ft ² /d)
32	Mrg-0148	140	45	6	635	44.09	11-4-2004	341CMNG	4.1	.64	54
33	Mrg-0120	160	25	6	840	4.27	8-25-2004	351TNLY	6.1	.55	45
34	Mrg-0116	195	--	6	--	2.36	8-12-2004	341CMNG	--	--	--
35	Mrg-0155	240	18.85	6	778	98.38	4-13-2005	341CMNG	5.4	.63	20
36	Mrg-0106	380	98	6	1,100	104.35	8-3-2004	341HMPR	--	.23	7
37	Mrg-0170	320	61.5	6	1,013	92.17	4-28-2005	341HMPR	3.2	.26	16
38	Mrg-0100	--	--	6	940	227.25	7-15-2004	341HMPR	--	--	--
39	Mrg-0085	219	--	6	830	--	5-19-2004	351TNLY	--	--	--
40	Mrg-0121	200	--	6	890	--	8-25-2004	354CLNN	--	--	--
41	Mrg-0160	150	38.45	6	1,129	60.01	4-19-2005	337POCN	5	0.13	5
42	Mrg-0123	205	--	6	920	34.3	8-26-2004	351TNLY	--	--	--
43	Mrg-0113	130	46	6	910	16.5	8-10-2004	354CLNN	--	3.72	266
44	Mrg-0088	560	--	6	980	--	7-1-2004	341HMPR	--	--	--
45	Mrg-0118	208	--	6	860	8.6	8-24-2004	341MRCL	--	.99	128
46	Mrg-0110	143	--	6	910	17.61	8-5-2004	351TNLY	--	.07	3
47	Mrg-0167	260	58	6	642	104.81	4-25-2005	351TNLY	4.3	1.02	34
48	Mrg-0097	300	60	6	920	31.95	7-13-2004	351TNLY	--	.06	2
49	Mrg-0117	260	135	6	880	95.75	6-29-2004	351TNLY	10.6	.63	21
50	Mrg-0133	240	60	6	910	27.75	9-9-2004	341CMNG	10.5	.47	55
51	Mrg-0171	205	38.75	6	913	84.33	5-2-2005	341CMNG	10	.25	9.5
52	Mrg-0172	160	18.0	6	869	34.52	5-2-2005	341CMNG	3.5	.15	4
53	Mrg-0173	62	8.0	6	787	8.47	5-3-2005	351TNLY	4.6	.51	25
54	Mrg-0103	123	25	6	930	25.5	7-28-2004	341CMNG	5.2	.28	11.5
55	Mrg-0102	305	60	6	930	22.15	7-28-2004	341CMNG	14.2	.44	48
56	Mrg-0101	310	62	6	820	7.97	7-27-2004	341BRLR	5.6	.27	21
57	Mrg-0105	320	--	6	890	154.73	8-3-2004	341CMNG	--	--	--
58	Mrg-0093	150	33	6	680	27.88	7-8-2004	341CMNG	5.4	.09	11
59	Mrg-0107	180	159	6	1,140	61.93	8-4-2004	354CLNN	--	.38	43
60	Mrg-0131	150	70	6	831	93.83	9-8-2004	341CMNG	3.1	.75	26
61	Mrg-0132	130	30	6	848	1	9-8-2004	351TNLY	1	.71	3
62	Mrg-0128	300	20	6	845	160.02	9-1-2004	341CMNG	4.7	.14	8

Appendix 1. Well-construction and aquifer properties for selected wells, Morgan County, West Virginia. — Continued

[USGS, U.S. Geological Survey; ft bls, feet below land surface; ft, foot; in., inch; Hydrogeologic unit codes: 354CLNN, Clinton Group; 351TNLY, Tomoloway Formation; 344MRCL, Marcellus Formation; 344MNNG, Mahantango Formation; 341BRLR, Brallier Formation; 341CMNG, Chemung Group; 341HMPR, Hampshire Formation; 337POCN, Pocomo Group, gal/min, gallons per minute; (gal/min)/ft, gallons per minute per foot of drawdown; ft²/d, foot squared per day; --, not determined. See figure 9 for well locations]

Site number	USGS station name	Well depth (ft bls)	Length of casing (ft)	Casing diameter (in.)	Land-surface altitude (ft)	Static water level (ft bls)	Date water level measured	Hydrogeologic unit	Discharge (gal/min)	Specific capacity ((gal/min)/ft)	Transmissivity average (ft ² /d)
63	Mrg-0125	420	79	6	723	63.2	8-31-2004	341BRLR	7.3	0.11	4.5
64	Mrg-0126	400	40	6	701	25.23	8-31-2004	341BRLR	6.2	.06	2
65	Mrg-0127	720	60	6	835	87.22	9-1-2004	341CMNG	4.3	.10	3
66	Mrg-0096	225	40	6	640	53.77	7-13-2004	341BRLR	6.5	.82	24
67	Mrg-0108	260	58	6	800	112.51	8-4-2004	341BRLR	--	.34	13
68	Mrg-0129	340	80	6	925	56.95	9-2-2004	341MRCL	4.5	.05	2
69	Mrg-0109	--	--	6	780	185.81	8-4-2004	341BRLR	--	--	--
70	Mrg-0119	300	--	6	700	3.55	8-24-2004	341BRLR	--	--	--
71	Mrg-0099	460	--	6	1,200	--	7-14-2004	341HMPR	7.1	--	--
72	Mrg-0098	405	--	6	1,230	146.77	7-14-2004	341HMPR	7.6	--	--
73	Mrg-0168	80	39.2	6	818	3	4-26-2005	344MNNG	4.3	2.40	327
74	Mrg-0111	400	19	6	740	4.65	8-9-2004	341BRLR	--	--	--
75	Mrg-0169	320	39	6	888	50.7	4-26-2005	341BRLR	3.2	.12	4
76	Mrg-0158	51	27.1	6	561	15.75	4-14-2005	341HMPR	.8	.15	40
77	Mrg-0089	180	40	6	900	36.85	7-1-2004	341CMNG	--	3.26	1,490
78	Mrg-0087	425	--	6	880	56.85	6-30-2004	341CMNG	10.7	1.06	76.5
79	Mrg-0166	480	--	6	800	268.45	4-22-2005	344MRCL	5.6	--	--
80	Mrg-0175	185	58.6	6	915	53.4	5-5-2005	341CMNG	7	.28	15
81	Mrg-0086	186	--	6	1,010	--	5-19-2004	341CMNG	--	--	--
82	Mrg-0154	121	21.8	6	848	32.9	4-12-2005	341HMPR	4.2	.52	38
83	Mrg-0164	300	38.9	6	866	64.63	4-21-2005	344MNNG	60	.89	195
84	Mrg-0174	440	--	6	1,003	281.33	5-3-2005	344MRCL	6	--	--
85	Mrg-0163	300	40.5	6	913	137.86	4-20-2005	344MNNG	7.2	6.32	944
86	Mrg-0165	400	--	6	853	226.77	4-21-2005	344MRCL	2.5	--	--
87	Mrg-0134	250	80	6	1,025	35.23	9-10-2004	341CMNG	7.9	.41	14
88	Mrg-0115	220	99	6	1,110	38.6	8-11-2004	341HMPR	--	.41	22
89	Mrg-0153	220	58.85	6	1,125	69.6	4-12-2005	341HMPR	4.8	.16	5
90	Mrg-0162	187	47.3	6	1,128	45.8	4-20-2005	341HMPR	2.8	.06	2
91	Mrg-0150	135	--	6	1,025	50.21	3-17-2005	341HMPR	3.2	--	--
92	Mrg-0157	103	--	6	656	19.70	4-14-2005	341BRLR	4.5	0.08	2

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0090 is the 90th well site established in Morgan County, West Virginia.

Appendix 2a. Water-quality data values for wells, Morgan County, West Virginia

[USGS, U.S. Geological Survey. Units of measure: deg C, degrees Celsius; µS/cm, microsiemens per centimeter at 25oC; mg/L, milligrams per liter; NTU, nephelometric turbidity units. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Temperature, (deg C)	Specific conductance (µS/cm)	pH (units)	Oxygen, dissolved (mg/L)	Turbidity (NTU)	Alkalinity, (mg/L as CaCO ₃)	Bicarbonate, dissolved, field (mg/L as HCO ₃)	Carbonate, dissolved, field (mg/L as CO ₃)	Water hardness (mg/L as CaCO ₃)	Calcium, dissolved (mg/L)
1	Mrg-0090	7-6-2004	13.5	160	5.7	5.6	<50	--	--	--	--	--
2	Mrg-0178	5-10-2005	12.8	164	6.8	2.6	55	67	82	<1	89	28.6
3	Mrg-0146	11-3-2004	12.8	230	6.4	E.4	19	72	88	<1	92	16.4
4	Mrg-0114	8-11-2004	12.6	220	6.8	E.3	10	106	129	<1	99	17.1
5	Mrg-0124	8-26-2004	13.4	226	6.7	--	<2.0	105	128	<1	100	22
6	Mrg-0104	8-2-2004	14.2	321	6.9	--	4.8	125	152	<1	150	31.7
7	Mrg-0122	8-25-2004	12.3	284	6.5	--	2.4	85	104	<1	130	19.9
8	Mrg-0144	11-2-2004	12.9	254	7.1	E.2	--	121	148	<1	--	E.01
9	Mrg-0130	9-2-2004	13.4	80	4.8	5.7	<1.0	8	10	<1	18	1.81
9	Mrg-0130	9-9-2004	13.6	77	4.7	4.8	<1.0	--	--	--	--	--
10	Mrg-0095	7-12-2004	13.2	209	6.4	E.6	<1.0	105	128	<1	82	15.6
11	Mrg-0139	10-19-2004	13.5	169	6.6	--	--	113	138	<1	66	13.9
12	Mrg-0136	9-22-2004	13.8	266	6.2	E.5	7	80	98	<1	100	17.5
13	Mrg-0142	10-21-2004	12.6	228	6.4	E.7	--	83	101	<1	91	18.1
14	Mrg-0145	10-19-2004	12.2	296	6.8	E.2	--	173	211	<1	130	38.9
15	Mrg-0135	9-22-2004	13.2	403	6.8	2.3	<1.0	131	160	<1	160	39.3
16	Mrg-0156	4-13-2005	12.1	157	6.5	2.0	<1.0	66	80	<1	63	17.7
17	Mrg-0094	7-12-2004	12.6	273	7.3	E.2	2.7	110	134	<1	62	15.3
18	Mrg-0161	4-19-2005	14.3	440	7.2	E.3	<1.0	184	224	<1	220	69.2
19	Mrg-0137	9-23-2004	14.0	201	6.8	E.4	100	76	93	<1	77	14.7
20	Mrg-0143	11-2-2004	12.9	178	6.4	1.7	<1.0	71	87	<1	69	15.7
21	Mrg-0140	10-20-2004	12.2	309	7.6	E.3	--	164	200	<1	140	40
22	Mrg-0092	7-7-2004	13.8	225	7.3	E.4	<3.0	106	129	<1	97	25.3
23	Mrg-0149	11-4-2004	11.6	207	6.8	E.2	<1.0	69	84	<1	78	18.3
24	Mrg-0147	11-3-2004	12.1	198	6.1	E.9	29	49	60	<1	90	18.4
25	Mrg-0112	8-10-2004	13.8	202	6.2	E.9	<1.0	77	94	<1	61	10.1
26	Mrg-0138	9-23-2004	13.0	227	6.6	E.3	1.7	65	79	<1	98	25
27	Mrg-0141	10-20-2004	12.9	283	6.3	E.4	--	69	84	<1	100	21.4
28	Mrg-0176	5-6-2005	12.5	213	6.9	E.8	4.7	93	113	<1	91	18.2
29	Mrg-0159	4-15-2005	12.4	458	5.0	E.8	5	10	12	<1	160	20.5

Appendix 2a. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: deg C, degrees Celsius; $\mu\text{S}/\text{cm}$, microsiemens per centimeter at 25°C; mg/L, milligrams per liter; NTU, nephelometric turbidity units. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Temperature, (deg C)	Specific conductance ($\mu\text{S}/\text{cm}$)	pH (units)	Oxygen, dissolved (mg/L)	Turbidity (NTU)	Alkalinity, (mg/L as CaCO_3)	Bicarbonate, dissolved, field (mg/L as HCO_3)	Carbonate, dissolved, field (mg/L as CO_3)	Water hardness (mg/L as CaCO_3)	Calcium, dissolved (mg/L)
30	Mrg-0091	7-7-2004	13.0	137	6.6	³ 1.4	2	64	78	<1	57	11.7
31	Mrg-0179	5-11-2005	14.1	298	6.7	E.3	20	99	121	<1	110	32.1
32	Mrg-0148	11-4-2004	12.2	88	5.9	5.3	<1.0	24	29	<1	32	6.14
33	Mrg-0120	8-25-2004	12.4	318	7.3	--	<1.0	163	199	<1	180	62
34	Mrg-0116	8-12-2004	12.9	368	6.9	E.4	4	185	226	<1	190	60.4
35	Mrg-0155	4-13-2005	12.5	322	7.2	E.3	2.4	138	168	<1	140	33.8
36	Mrg-0106	8-3-2004	13.5	125	6.0	--	<10	64	78	<1	51	14
37	Mrg-0170	4-28-2005	12.5	256	7.0	E.9	<1.0	17	130	<1	110	23.4
38	Mrg-0100	7-15-2004	12.6	293	7.7	E.2	<1.0	127	155	<1	94	23.1
39	Mrg-0085	5-19-2004	13.8	502	7.3	³ 13	<1.0	176	214	<1	--	--
40	Mrg-0121	8-25-2004	12.4	158	6.7	--	<1.0	69	84	<1	80	27.6
41	Mrg-0160	4-19-2005	12.2	140	6.2	1.1	3.8	67	82	<1	58	10.2
42	Mrg-0123	8-26-2004	12.8	737	6.8	2.6	<15	305	372	<1	400	127
43	Mrg-0113	8-10-2004	11.8	160	6.1	8.7	25	75	91	<1	76	23.2
44	Mrg-0088	7-1-2004	12.4	245	7.3	E.4	<1.0	129	157	<1	100	20.7
45	Mrg-0118	8-24-2004	15.1	214	6.1	E.3	5	71	87	<1	87	25.9
46	Mrg-0110	8-5-2004	13.0	456	7.1	2.3	91	210	256	<1	270	73.8
47	Mrg-0167	4-25-2005	11.7	323	7.4	⁴ 5.9	24	141	172	<1	160	45.8
48	Mrg-0097	7-13-2004	14.2	393	7.0	4.3	<2.0	202	246	<1	210	56.2
49	Mrg-0117	6-29-2004	12.3	500	6.7	5.9	<5.0	209	255	<1	240	75.1
50	Mrg-0133	9-9-2004	13.8	478	7.0	E.3	22	238	290	<1	190	39
51	Mrg-0171	5-2-2005	12.2	272	7.3	E.4	2.7	119	145	<1	110	31.7
52	Mrg-0172	5-2-2005	11.8	227	6.5	E.4	100	94	115	<1	110	36.1
53	Mrg-0173	5-3-2005	11.2	402	7.2	E.3	17	189	230	<1	220	76.3
54	Mrg-0103	7-28-2004	14.7	279	6.3	--	9.1	159	194	<1	130	23.2
55	Mrg-0102	7-28-2004	13.1	242	6.3	E.7	<1.0	114	139	<1	90	13.1
56	Mrg-0101	7-27-2004	13.2	202	6.7	E.6	<1.0	110	134	<1	95	20.8
57	Mrg-0105	8-3-2004	14.1	158	6.0	--	310	30	37	<1	71	10.8
58	Mrg-0093	7-8-2004	13.9	180	6.4	--	--	80	98	<1	78	20.8
59	Mrg-0107	8-4-2004	13.1	198	6.3	6	<20	70	85	<1	110	39.8

Appendix 2a. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: deg C, degrees Celsius; µS/cm, microsiemens per centimeter at 25oC; mg/L, milligrams per liter; NTU, nephelometric turbidity units. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Temperature, (deg C)	Specific conductance (µS/cm)	pH (units)	Oxygen, dissolved (mg/L)	Turbidity (NTU)	Alkalinity, (mg/L as CaCO ₃)	Bicarbonate, dissolved, field (mg/L as HCO ₃)	Carbonate, dissolved, field (mg/L as CO ₃)	Water hardness (mg/L as CaCO ₃)	Calcium, dissolved (mg/L)
60	Mrg-0131	9-8-2004	13.4	182	6.5	4.7	1.3	129	157	<1	82	13.7
61	Mrg-0132	9-8-2004	16.0	451	7.3	1.4	8	231	282	<1	260	53.7
62	Mrg-0128	9-1-2004	13.4	305	7.4	2.1	25	171	208	<1	82	20.2
63	Mrg-0125	8-31-2004	13.2	294	7.4	E.5	4	133	162	<1	130	29.9
64	Mrg-0126	8-31-2004	14.1	269	7.6	E.6	23	107	130	<1	89	25.1
65	Mrg-0127	9-1-2004	13.5	167	6.3	1	<1.0	84	102	<1	70	10.2
66	Mrg-0096	7-13-2004	13.2	249	6.6	E.8	4.8	94	115	<1	110	33.9
67	Mrg-0108	8-4-2004	12.4	275	7.6	--	1.2	130	158	<1	130	33.3
68	Mrg-0129	9-2-2004	13.7	313	6.7	E.3	59	140	171	<1	150	51.5
69	Mrg-0109	8-4-2004	14.2	271	7.2	--	1.5	107	130	<1	120	28.8
70	Mrg-0119	8-24-2004	13.1	270	7.4	--	<20	96	117	<1	120	33.8
71	Mrg-0099	7-14-2004	13.5	252	7.6	1.9	<1.0	115	140	<1	110	23.8
72	Mrg-0098	7-14-2004	13.3	263	7.1	2.9	15	117	143	<1	130	25.3
73	Mrg-0168	4-26-2005	12.3	435	6.6	E.6	5.7	83	101	<1	160	34.8
74	Mrg-0111	8-9-2004	13.9	270	7.2	<2	<10	127	155	<1	120	28.5
75	Mrg-0169	4-26-2005	12.7	322	6.7	E.5	1.2	111	135	<1	140	37
76	Mrg-0158	4-14-2005	12.4	222	6.9	E.3	11	94	115	<1	87	22.3
77	Mrg-0089	7-1-2004	12.5	224	7.6	E.3	<2.0	108	132	<1	98	27.2
78	Mrg-0087	6-30-2004	13.4	235	7.3	E.1	<2.0	113	138	<1	100	25.5
79	Mrg-0166	4-22-2005	13.8	75	5.0	8.9	<1.0	8	10	<1	27	7.75
80	Mrg-0175	5-5-2005	13.2	195	6.6	E.4	4.3	83	101	<1	86	14.2
81	Mrg-0086	5-19-2004	13.8	252	6.3	E.8	2.4	98	119	<1	--	212
82	Mrg-0154	4-12-2005	11.9	44	5.3	5.1	2	17	21	<1	16	2.15
83	Mrg-0164	4-21-2005	12.4	239	6.6	2.4	1.1	94	115	<1	98	19.7
84	Mrg-0174	5-3-2005	12.6	82	5.1	10.3	330	11	13	<1	30	8.82
85	Mrg-0163	4-20-2005	13	227	6.9	E.5	<1.0	84	102	<1	98	28.9
86	Mrg-0165	4-21-2005	13.5	103	5.4	5.4	10	22	27	<1	38	10.4
87	Mrg-0134	9-10-2004	12.6	241	6.4	E.6	17	110	134	<1	100	17.1
88	Mrg-0115	8-11-2004	12.8	263	7.4	<2	<4.0	137	167	<1	130	25.5
89	Mrg-0153	4-12-2005	12	124	6.4	2.5	1.8	47	57	<1	50	8.77

Appendix 2a. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: deg C, degrees Celsius; $\mu\text{S}/\text{cm}$, microsiemens per centimeter at 25°C; mg/L, milligrams per liter; NTU, nephelometric turbidity units. Symbols: --, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Temperature (deg C)	Specific conductance ($\mu\text{S}/\text{cm}$)	pH (units)	Oxygen dissolved (mg/L)	Turbidity (NTU)	Alkalinity (mg/L as CaCO_3)	Bicarbonate dissolved, field (mg/L as HCO_3)	Carbonate dissolved, field (mg/L as CO_3)	Water hardness (mg/L as CaCO_3)	Calcium, dissolved (mg/L)
90	Mrg-0162	4-20-2005	12	131	6.8	2	3.8	57	69	<1	58	12.7
91	Mrg-0150	3-17-2005	10.9	163	7.0	0.7	51	79	96	<1	63	13

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0090 is the 90th well site established in Morgan County, West Virginia.

²Sample concentration reported as "Total."

³Water cascading down well bore during test.

⁴Air entrained in plumbing system.

Appendix 2b. Water-quality data values for wells, Morgan County, West Virginia

[USGS, U.S. Geological Survey. Units of measure: mg/L, milligrams per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Magnesium dissolved (mg/L)	Potassium dissolved (mg/L)	Sodium dissolved (mg/L)	Bromide dissolved (mg/L)	Chloride dissolved (mg/L)	Fluoride dissolved (mg/L)	Silica dissolved (mg/L)	Sulfate dissolved (mg/L)	Solids, residue at 180 deg C dissolved (mg/L)	Ammonia, dissolved (mg/L as NH4)	Ammonia, dissolved (mg/L as N)
1	Mrg-0090	7-6-2004	--	--	--	--	--	--	--	--	--	--	--
2	Mrg-0178	5-10-2005	4.16	0.83	0.72	--	3.23	0.2	8.3	23.5	115	--	<0.04
3	Mrg-0146	11-3-2004	12.5	.67	7.69	--	21.9	.2	21.2	3.8	131	--	<.04
4	Mrg-0114	8-11-2004	13.7	.71	9.78	--	4.03	.2	22.8	3.5	136	--	E.02
5	Mrg-0124	8-26-2004	11.8	.59	11.9	--	1.17	<.2	24.7	3.1	141	0.08	.06
6	Mrg-0104	8-2-2004	16.7	.59	8.38	--	24.7	.2	20.3	5.1	183	.05	.04
7	Mrg-0122	8-25-2004	18.7	.63	9.14	--	5.04	.2	22.1	52.6	185	--	E.02
8	Mrg-0144	11-2-2004	<.008	<.16	59	--	.92	.1	13.1	2.8	--	--	<.04
9	Mrg-0130	9-2-2004	3.31	.44	5.28	0.12	14.3	<.2	8	3.1	43	--	<.04
9	Mrg-0130	9-9-2004	--	--	--	--	--	--	--	--	--	--	--
10	Mrg-0095	7-12-2004	10.4	.71	9.68	--	2.08	.2	22.2	2.8	130	--	<.04
11	Mrg-0139	10-19-2004	7.53	.64	8.05	--	1.16	.2	23	2.6	126	--	<.04
12	Mrg-0136	9-22-2004	14.9	.79	8.82	--	27.4	.2	28.5	2.9	151	.06	.05
13	Mrg-0142	10-21-2004	11.1	.7	9.76	--	13.8	.2	29.8	2.6	140	--	E.03
14	Mrg-0145	10-19-2004	8.68	.48	10.6	.02	5.16	.2	35.1	3.6	209	--	E.03
15	Mrg-0135	9-22-2004	13.8	.58	18.3	--	28.5	<.2	15.2	10.7	206	--	<.04
16	Mrg-0156	4-13-2005	4.5	.57	7.16	--	1.09	.1	19	9.3	99	--	<.04
17	Mrg-0094	7-12-2004	5.68	.39	33.1	--	6.56	.2	14.7	8.7	151	.06	.05
18	Mrg-0161	4-19-2005	12.5	1.82	3.63	--	5.92	.1	8.1	34.3	246	.22	.17
19	Mrg-0137	9-23-2004	9.75	.82	9.95	--	7.22	.2	17.4	5.7	112	--	<.04
20	Mrg-0143	11-2-2004	7.26	.71	9.03	--	1.4	.1	21.7	12.4	111	--	<.04
21	Mrg-0140	10-20-2004	8.75	.37	11.8	--	2.24	.1	19	8.8	190	--	<.04
22	Mrg-0092	7-7-2004	8.25	.34	8.16	.02	2.17	.2	17.7	7.5	134	--	<.04
23	Mrg-0149	11-4-2004	7.78	.71	8.87	--	10.3	.2	20.6	9	118	--	<.04
24	Mrg-0147	11-3-2004	10.7	.92	5.14	--	.86	E.1	27.1	36.6	130	--	<.04
25	Mrg-0112	8-10-2004	8.6	.72	10.1	--	17.2	<.2	27.5	1.3	129	.15	.12
26	Mrg-0138	9-23-2004	8.68	.67	5.83	--	2.35	<.2	28.1	30.5	143	--	E.02
27	Mrg-0141	10-20-2004	11.4	1.04	13.1	--	36.2	.1	26.9	6.7	160	.07	.05
28	Mrg-0176	5-6-2005	11.1	.7	9.38	--	1.36	.3	22.7	8	128	--	E.03
29	Mrg-0159	4-15-2005	25.9	2.09	1.16	--	.76	<.1	11.6	172	268	.07	.06
30	Mrg-0091	7-7-2004	6.69	.71	5.41	--	.88	<.2	18.1	5.2	88	--	<.04
31	Mrg-0179	5-11-2005	6.43	.74	.78	--	.78	.3	21.1	6.1	128	--	<.04

Appendix 2b. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: mg/L, milligrams per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Magnesium, dissolved (mg/L)	Potassium, dissolved (mg/L)	Sodium, dissolved (mg/L)	Bromide, dissolved (mg/L)	Chloride, dissolved (mg/L)	Fluoride, dissolved (mg/L)	Silica, dissolved (mg/L)	Sulfate, dissolved (mg/L)	Solids, residue at 180 deg C dissolved (mg/L)	Ammonia, dissolved (mg/L as NH4)	Ammonia, dissolved (mg/L as N)
32	Mrg-0148	11-4-2004	4.01	0.61	4.04	--	2.52	E0.1	14.9	4.2	57	--	<.04
33	Mrg-0120	8-25-2004	6.25	.7	2.24	--	1.51	<.2	8.1	11	190	0.06	.05
34	Mrg-0116	8-12-2004	8.59	.58	7.85	0.05	5.15	<.2	21.4	8.9	226	.06	.05
35	Mrg-0155	4-13-2005	13.5	.46	9.59	--	7.32	.2	22.4	4.3	175	.1	.08
36	Mrg-0106	8-3-2004	3.93	.78	6.36	--	.89	<.2	23.2	5.7	94	--	<.04
37	Mrg-0170	4-28-2005	13.2	1.31	9.19	--	3.4	.2	19	6.7	141	--	<.04
38	Mrg-0100	7-15-2004	8.8	1.23	26.4	--	2.92	<.2	14.1	12.1	165	--	<.04
39	Mrg-0085	5-19-2004	23	8	35.7	<.05	17	2<.1	--	27	--	2.03	2.02
40	Mrg-0121	8-25-2004	2.77	.5	1.27	--	.96	<.2	7.2	4	86	--	<.04
41	Mrg-0160	4-19-2005	7.8	.93	7.99	--	.9	.2	22.8	2.5	94	--	<.04
42	Mrg-0123	8-26-2004	19.8	.89	13.9	.13	8.8	.2	8.6	71.3	444	--	<.04
43	Mrg-0113	8-10-2004	4.41	.74	2.07	.03	2.95	<.2	7.2	3.9	91	--	<.04
44	Mrg-0088	7-1-2004	11.9	1.22	14.5	--	1.32	<.2	15.6	4.8	147	--	<.04
45	Mrg-0118	8-24-2004	5.32	.72	3.49	.02	1.17	<.2	21.9	30.4	139	--	E.04
46	Mrg-0110	8-5-2004	21.5	.53	.85	--	2.13	<.2	6.7	5.9	238	--	<.04
47	Mrg-0167	4-25-2005	12.2	.83	.79	--	.96	.1	6.9	7.7	160	--	<.04
48	Mrg-0097	7-13-2004	16.5	1.1	.83	--	1.72	<.2	7.1	4	208	--	<.04
49	Mrg-0117	6-29-2004	12.1	.73	15	--	32.5	<.2	7.7	3.5	275	--	<.04
50	Mrg-0133	9-9-2004	22.7	.42	29.7	--	24.7	.2	21.4	7.7	290	.07	.05
51	Mrg-0171	5-2-2005	8.07	.4	14.1	--	2.29	.2	17.2	8.1	153	--	<.04
52	Mrg-0172	5-2-2005	4.2	.37	3.53	--	1.53	E.1	17.2	7.8	127	--	<.04
53	Mrg-0173	5-3-2005	7.66	.39	2.13	--	1.96	.1	8.5	8.5	219	--	<.04
54	Mrg-0103	7-28-2004	16.6	.65	11.4	--	8.66	<.2	27.2	7.2	192	--	E.02
55	Mrg-0102	7-28-2004	14	.7	11.1	--	4.5	<.2	26	4	148	--	E.02
56	Mrg-0101	7-27-2004	10.5	.85	9.13	--	.67	.2	20.2	7	135	--	E.03
57	Mrg-0105	8-3-2004	10.7	.66	8.34	--	1.18	<.2	21.2	2.9	74	--	<.04
58	Mrg-0093	7-8-2004	6.33	.41	6.97	.04	6.12	<.2	20.4	7	118	--	<.04
59	Mrg-0107	8-4-2004	1.74	.47	.49	--	.55	<.2	7.8	6.6	99	--	<.04
60	Mrg-0131	9-8-2004	11.6	.53	9.34	.03	.81	.2	24.2	1.1	139	--	<.04
61	Mrg-0132	9-8-2004	29.4	.73	2.82	.04	3.23	<.2	11.8	15.2	269	--	<.04
62	Mrg-0128	9-1-2004	7.71	.36	38.6	--	.99	<.2	16.8	10.7	198	.08	.06
63	Mrg-0125	8-31-2004	12.6	.31	15.3	--	1.08	.2	17	12.2	168	.15	.12

Appendix 2b. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: mg/L, milligrams per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Magnesium, dissolved (mg/L)	Potassium, dissolved (mg/L)	Sodium, dissolved (mg/L)	Bromide, dissolved (mg/L)	Chloride, dissolved (mg/L)	Fluoride, dissolved (mg/L)	Silica, dissolved (mg/L)	Sulfate, dissolved (mg/L)	Solids,		
											residue at 180 deg C dissolved (mg/L)	Ammonia, dissolved (mg/L as NH4)	Ammonia, dissolved (mg/L as N)
64	Mrg-0126	8-31-2004	6.5	0.43	22.7	--	2.29	0.2	16.1	21	159	0.2	0.15
65	Mrg-0127	9-1-2004	10.8	.63	8.68	--	.87	<2	24.2	2.4	110	--	E.03
66	Mrg-0096	7-13-2004	5.24	.7	7.26	--	1.86	<2	19.5	9.1	134	--	<.04
67	Mrg-0108	8-4-2004	11.4	.39	10.2	--	1.37	<2	15.9	12.9	163	.07	.05
68	Mrg-0129	9-2-2004	4.73	.64	5.05	--	11	<2	9.7	17.2	184	--	<.04
69	Mrg-0109	8-4-2004	11.8	.34	12.8	--	1.85	.2	16.5	22.6	159	--	<.04
70	Mrg-0119	8-24-2004	7.16	.63	15.6	0.02	1.89	.2	15.8	18.9	153	.22	.17
71	Mrg-0099	7-14-2004	11.9	1.18	11.7	--	1.54	<2	14.8	5.8	140	--	<.04
72	Mrg-0098	7-14-2004	15.6	1.57	6.78	--	1.47	<2	14.2	8.9	151	--	<.04
73	Mrg-0168	4-26-2005	18.8	.41	15.3	--	60.7	.1	27.8	21	232	.26	.2
74	Mrg-0111	8-9-2004	11.9	.37	13.3	.03	1.66	.2	15.7	17.8	166	--	<.04
75	Mrg-0169	4-26-2005	11.2	.77	10.8	--	3.77	0.3	29.6	29.8	192	--	E.03
76	Mrg-0158	4-14-2005	7.57	.5	9.12	--	1.84	.2	19.9	8.8	129	.09	.07
77	Mrg-0089	7-1-2004	7.21	.33	8.53	--	1.2	<2	15.2	7.9	133	--	E.04
78	Mrg-0087	6-30-2004	9.27	.47	9.41	--	3.66	.2	16	5.6	138	--	<.04
79	Mrg-0166	4-22-2005	1.87	1.19	.69	--	1.45	<.1	6.9	13.4	46	--	<.04
80	Mrg-0175	5-5-2005	12.2	.79	7.81	--	6.06	.2	22.9	2.1	117	--	<.04
81	Mrg-0086	5-19-2004	² 15	² .7	² 12	.12	6.6	² .2	--	1.4	--	² <.01	² <.02
82	Mrg-0154	4-12-2005	2.64	.46	3.4	--	.68	<.1	15.2	1.8	37	--	<.04
83	Mrg-0164	4-21-2005	11.7	.66	11.7	--	.65	.2	22.8	17.9	143	.08	.06
84	Mrg-0174	5-3-2005	1.82	1.7	.79	--	1.75	<.1	6.6	11.8	40	--	<.04
85	Mrg-0163	4-20-2005	6.32	.72	7.77	--	.97	.2	19.7	17.3	132	--	<.04
86	Mrg-0165	4-21-2005	2.89	.83	3.51	--	6.43	<.1	8.1	12.4	58	--	<.04
87	Mrg-0134	9-10-2004	15.1	.49	10.8	--	5.52	<.2	29.1	3	148	--	E.04
88	Mrg-0115	8-11-2004	14.9	1.42	9.79	--	1.83	<.2	16.7	4.3	157	--	<.04
89	Mrg-0153	4-12-2005	6.74	.74	5.11	--	.75	.1	20.1	4.7	76	--	<.04
90	Mrg-0162	4-20-2005	6.33	.7	4.25	--	.67	.1	16.3	5.2	81	--	<.04
91	Mrg-0150	3-17-2005	7.47	.48	9.02	--	.96	.1	19.4	4.7	103	--	<.04

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0090 is the 90th well site established in Morgan County, West Virginia.²Sample concentration reported as "Total."

Appendix 2c. Water-quality data values for wells, Morgan County, West Virginia.

[USGS, U.S. Geological Survey, Units of measure: mg/L, milligrams per liter; col/100 ml, colonies per 100 milliliters; µg/L, micrograms per liter. Symbols:--, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Nitrate+ nitrite, dissolved (mg/L as N)	Nitrite, dissolved (mg/L as N)	Ortho-phosphate, dissolved (mg/L)	Ortho-phosphate, dissolved (mg/L as P)	Phosphorus, dissolved (mg/L)	<i>E. coli</i> , (col/100 ml)	Coliform, fecal (col/100 ml)	Coliform, total (col/100 ml)	Aluminum, total (µg/L)	Antimony, dissolved (µg/L)
1	Mrg-0090	7-6-2004	--	--	--	--	--	--	--	--	--	--
2	Mrg-0178	5-10-2005	1.11	<0.008	<0.02	<0.04	<0.04	<1	<1	<1	--	--
3	Mrg-0146	11-3-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	E4	--	--
4	Mrg-0114	8-11-2004	<.06	<.008	E.01	E.03	E.03	<1	<1	<1	--	--
5	Mrg-0124	8-26-2004	<.06	<.008	<.02	E.04	E.04	<1	<1	<1	--	--
6	Mrg-0104	8-2-2004	<.06	<.008	E.01	<.04	<.04	<1	<1	E12	--	--
7	Mrg-0122	8-25-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	<1	--	--
8	Mrg-0144	11-2-2004	<.06	<.008	.03	.04	.04	<1	<1	<1	--	--
9	Mrg-0130	9-2-2004	.34	<.008	<.02	<.04	<.04	--	--	--	70	<0.20
9	Mrg-0130	9-9-2004	--	--	--	--	--	<1	<1	E17	--	--
10	Mrg-0095	7-12-2004	<.06	<.008	<.02	<.04	<.04	--	<1	--	--	--
11	Mrg-0139	10-19-2004	<.06	<.008	<.02	E.03	E.03	<1	<1	E2	--	--
12	Mrg-0136	9-22-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	E7	--	--
13	Mrg-0142	10-21-2004	<.06	<.008	<.02	<.09	.09	<1	<1	E380	--	--
14	Mrg-0145	10-19-2004	<.06	<.008	<.02	E.02	E.02	<1	<1	<1	13	--
15	Mrg-0135	9-22-2004	.18	<.008	<.02	<.04	<.04	<1	<1	<1	--	--
16	Mrg-0156	4-13-2005	.08	<.008	<.02	<.04	<.04	<1	<1	<1	--	--
17	Mrg-0094	7-12-2004	<.06	<.008	.095	<.03	E.04	<1	<1	<1	--	--
18	Mrg-0161	4-19-2005	<.06	<.008	<.02	<.04	<.04	<1	<1	<1	<2	--
19	Mrg-0137	9-23-2004	<.06	<.008	E.01	E.03	E.03	<1	E38	22	--	--
20	Mrg-0143	11-2-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	120	--	--
21	Mrg-0140	10-20-2004	E.04	<.008	<.02	<.04	<.04	<1	<1	E5	--	--
22	Mrg-0092	7-7-2004	E.04	.042	<.02	<.04	<.04	<1	<1	70	50	--
23	Mrg-0149	11-4-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	E12	--	--
24	Mrg-0147	11-3-2004	<.06	<.008	E.02	<.04	<.04	<1	<1	E5	--	--
25	Mrg-0112	8-10-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	<1	17	<.20
26	Mrg-0138	9-23-2004	<.06	<.008	<.02	<.04	<.04	<1	350	100	--	--
27	Mrg-0141	10-20-2004	<.06	<.008	<.02	<.04	<.04	<1	<1	E4	--	--
28	Mrg-0176	5-6-2005	<.06	<.008	E.01	<.04	<.04	<1	<1	49	7	E.13
29	Mrg-0159	4-15-2005	<.06	<.008	<.02	<.04	<.04	<1	<1	E5	--	--
30	Mrg-0091	7-7-2004	.09	<.008	.06	.07	.07	<1	<1	E13	--	--
31	Mrg-0179	5-11-2005	<.06	<.008	E.01	<.04	<.04	<1	--	<1	--	--

Appendix 2c. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey, Units of measure: mg/L, milligrams per liter; col/100 ml, colonies per 100 milliliters; µg/L, micrograms per liter. Symbols:—, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Nitrate+ nitrite, dissolved (mg/L as N)	Nitrite, dissolved (mg/L as N)	Ortho-phosphate, dissolved (mg/L)	Ortho-phosphate, dissolved (mg/L as P)	Phosphorus, dissolved (mg/L)	E. coli, (col/100 ml)	Coliform, fecal (col/100 ml)	Coliform, total (col/100 ml)	Aluminum, total (µg/L)	Antimony, dissolved (µg/L)
32	Mrg-0148	11-4-2004	1.51	<0.008	—	E0.02	E0.03	<1	<1	E4	—	—
33	Mrg-0120	8-25-2004	<.06	<.008	—	<.02	<.04	<1	<1	<1	—	—
34	Mrg-0116	8-12-2004	<.06	<.008	—	<.02	<.04	E2	<1	33	<2	<.20
35	Mrg-0155	4-13-2005	<.06	<.008	0.067	.02	E.02	<1	<1	<1	5	<.20
36	Mrg-0106	8-3-2004	E.04	<.008	.509	.17	.18	<1	<1	<1	—	—
37	Mrg-0170	4-28-2005	.12	<.008	.233	.08	.09	<1	<1	<1	—	—
38	Mrg-0100	7-15-2004	<.06	<.008	.107	.04	E.03	<1	<1	E3	—	—
39	Mrg-0085	5-19-2004	—	2<.01	—	—	—	—	—	—	<3	2<1
40	Mrg-0121	8-25-2004	.13	<.008	—	<.02	<.04	<1	<1	<1	—	—
41	Mrg-0160	4-19-2005	E.05	<.008	.31	.1	.1	<1	<1	<1	3	—
42	Mrg-0123	8-26-2004	2.06	<.008	—	<.02	<.04	<1	<1	<1	90	<.20
43	Mrg-0113	8-10-2004	.24	<.008	—	<.02	<.04	E3	<1	E13	110	<.20
44	Mrg-0088	7-1-2004	<.06	<.008	.209	.07	.07	<1	<1	24	—	—
45	Mrg-0118	8-24-2004	<.06	E.005	—	E.01	.11	<1	<1	<1	18	<.20
46	Mrg-0110	8-5-2004	.07	<.008	—	<.02	<.04	<1	E2	E4	—	—
47	Mrg-0167	4-25-2005	E.06	<.008	—	<.02	<.04	<1	<1	<1	2	—
48	Mrg-0097	7-13-2004	E.04	<.008	—	<.02	<.04	<1	<1	E1	—	—
49	Mrg-0117	6-29-2004	.62	<.008	—	<.02	<.04	<1	<1	E2	—	—
50	Mrg-0133	9-9-2004	<.06	<.008	—	<.02	<.04	<1	<1	<1	—	—
51	Mrg-0171	5-2-2005	<.06	<.008	—	<.02	<.04	<1	<1	52	—	—
52	Mrg-0172	5-2-2005	<.06	<.008	—	<.02	<.04	<1	<1	<1	—	—
53	Mrg-0173	5-3-2005	<.06	<.008	—	<.02	<.04	<1	<1	<1	13	<.20
54	Mrg-0103	7-28-2004	E.06	<.008	—	<.02	<.04	<1	<1	120	—	—
55	Mrg-0102	7-28-2004	<.06	E.004	—	<.02	E.02	<1	<1	480	—	—
56	Mrg-0101	7-27-2004	<.06	<.008	—	E.01	<.04	<1	<1	E3	—	—
57	Mrg-0105	8-3-2004	<.06	<.008	—	<.02	.25	<1	E2	<1	—	—
58	Mrg-0093	7-8-2004	.24	E.006	—	<.02	<.04	<1	<1	E1	E1	—
59	Mrg-0107	8-4-2004	<.06	<.008	—	<.02	<.04	<1	<1	<1	—	—
60	Mrg-0131	9-8-2004	<.06	<.008	.08	.03	E.04	<1	<1	<1	10	<.20
61	Mrg-0132	9-8-2004	1.83	<.008	—	<.02	<.04	<1	<1	<1	5	<.20
62	Mrg-0128	9-1-2004	<.06	<.008	—	<.02	<.04	<1	<1	E6	—	—
63	Mrg-0125	8-31-2004	<.06	<.008	—	E.01	<.04	<1	<1	E15	—	—

Appendix 2c. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey, Units of measure: mg/L, milligrams per liter; col/100 ml, colonies per 100 milliliters; µg/L, micrograms per liter. Symbols:--, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Nitrate+ nitrite, dissolved (mg/L as N)	Nitrite, dissolved (mg/L as N)	Ortho-phosphate, dissolved (mg/L)	Ortho-phosphate, dissolved (mg/L as P)	Phosphorus, dissolved (mg/L)	<i>E. coli</i> , (col/100 ml)	Coliform, fecal (col/100 ml)	Coliform, total (col/100 ml)	Aluminum, total (µg/L)	Antimony, dissolved (µg/L)
64	Mrg-0126	8-31-2004	<0.06	<0.008	--	<0.02	<0.04	<1	<1	<1	--	--
65	Mrg-0127	9-1-2004	<0.06	<0.008	--	<0.02	<0.04	<1	<1	<1	--	--
66	Mrg-0096	7-13-2004	<0.06	<0.008	--	E.01	E.03	<1	E1	110	--	--
67	Mrg-0108	8-4-2004	<0.06	<0.008	--	<0.02	<0.04	<1	<1	<1	--	--
68	Mrg-0129	9-2-2004	.07	<0.008	--	<0.02	<0.04	<1	<1	25	--	--
69	Mrg-0109	8-4-2004	.08	<0.008	--	<0.02	<0.04	<1	<1	<1	--	--
70	Mrg-0119	8-24-2004	<0.06	<0.008	--	<0.02	<0.04	<1	<1	<1	890	<0.20
71	Mrg-0099	7-14-2004	.08	<0.008	0.11	.04	E.03	<1	<1	<1	--	--
72	Mrg-0098	7-14-2004	1.42	<0.008	.169	.06	.06	<1	<1	E13	--	--
73	Mrg-0168	4-26-2005	<0.06	<0.008	.071	.02	.07	<1	<1	<1	3	--
74	Mrg-0111	8-9-2004	<0.06	<0.008	--	<0.02	.15	<1	<1	E12	--	--
75	Mrg-0169	4-26-2005	<0.06	<0.008	.071	.02	E.04	<1	<1	<1	7	--
76	Mrg-0158	4-14-2005	<0.06	<0.008	--	<0.02	<0.04	<1	<1	<1	--	--
77	Mrg-0089	7-1-2004	<0.06	<0.008	--	<0.02	<0.04	<1	<1	740	--	--
78	Mrg-0087	6-30-2004	<0.06	<0.008	--	<0.02	<0.04	<1	<1	30	--	--
79	Mrg-0166	4-22-2005	1.7	<0.008	.092	.03	.05	<1	<1	<1	12	--
80	Mrg-0175	5-5-2005	<0.06	<0.008	--	<0.02	<0.04	<1	<1	E8	8	<.20
81	Mrg-0086	5-19-2004	--	² <0.01	--	--	² .06	<1	<1	<1	<3	² <1
82	Mrg-0154	4-12-2005	<0.06	<0.008	--	E.01	<0.04	<1	<1	<1	50	<.20
83	Mrg-0164	4-21-2005	<0.06	<0.008	--	<0.02	<0.04	<1	<1	E2	--	--
84	Mrg-0174	5-3-2005	2.6	<0.008	.095	.03	E.03	<1	<1	<1	2,340	<.20
85	Mrg-0163	4-20-2005	<0.06	<0.008	--	E.01	<0.04	<1	<1	<1	2	--
86	Mrg-0165	4-21-2005	.15	<0.008	--	E.01	E.02	<1	<1	<1	290	--
87	Mrg-0134	9-10-2004	<0.06	<0.008	--	E.01	.09	<1	<1	650	--	--
88	Mrg-0115	8-11-2004	<0.06	<0.008	--	E.02	E.02	<1	<1	<1	--	--
89	Mrg-0153	4-12-2005	.16	<0.008	.279	.09	.09	<1	<1	<1	--	--
90	Mrg-0162	4-20-2005	.13	<0.008	.058	.02	E.03	<1	<1	<1	60	--
91	Mrg-0150	3-17-2005	E.05	<0.008	--	<0.02	<0.04	--	<1	--	--	--

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0090 is the 90th well site established in Morgan County, West Virginia.² Sample concentration reported as "Total."

Appendix 2d. Water-quality data values for wells, Morgan County, West Virginia.

[USGS, U.S. Geological Survey. Units of measure: µg/L, micrograms per liter; pCi/L, picocuries per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Arsenic, total (µg/L)	Barium, total (µg/L)	Cadmium, dissolved (µg/L)	Iron, dissolved (µg/L)	Lead, dissolved (µg/L)	Manganese, dissolved (µg/L)	Nickel, dissolved (µg/L)	Strontium, dissolved (µg/L)	Zinc, total (µg/L)	Radon-222, total (pCi/L)
1	Mrg-0090	7-6-2004	--	--	--	--	--	--	--	--	--	--
2	Mrg-0178	5-10-2005	--	--	--	<6	--	<0.6	--	--	--	--
3	Mrg-0146	11-3-2004	--	--	--	2,870	--	1,020	--	--	--	--
4	Mrg-0114	8-11-2004	--	--	--	375	--	394	--	--	--	--
5	Mrg-0124	8-26-2004	--	--	--	1,970	--	369	--	--	--	--
6	Mrg-0104	8-2-2004	--	--	--	298	--	449	--	--	--	--
7	Mrg-0122	8-25-2004	--	--	--	4,400	--	977	--	--	--	--
8	Mrg-0144	11-2-2004	--	--	--	7	--	<6	--	--	--	--
9	Mrg-0130	9-2-2004	<2	52.4	E0.04	<6	1.7	38.2	12.5	8.87	18	2,170
9	Mrg-0130	9-9-2004	--	--	--	--	--	--	--	--	--	--
10	Mrg-0095	7-12-2004	--	--	--	2,440	--	482	--	--	--	--
11	Mrg-0139	10-19-2004	--	--	--	944	--	298	--	--	--	--
12	Mrg-0136	9-22-2004	--	--	--	769	--	705	--	--	--	--
13	Mrg-0142	10-21-2004	--	--	--	2,950	--	965	--	--	--	--
14	Mrg-0145	10-19-2004	<2	344	--	2,150	--	698	--	--	E5	120
15	Mrg-0135	9-22-2004	--	--	--	E4	--	14	--	--	--	--
16	Mrg-0156	4-13-2005	--	--	--	<6	--	7.1	--	--	--	--
17	Mrg-0094	7-12-2004	--	--	--	E5	--	73.8	--	--	--	--
18	Mrg-0161	4-19-2005	<2	60.4	--	223	--	57.9	--	--	238	--
19	Mrg-0137	9-23-2004	--	--	--	9	--	526	--	--	--	--
20	Mrg-0143	11-2-2004	--	--	--	27	--	53.1	--	--	--	--
21	Mrg-0140	10-20-2004	--	--	--	13	--	102	--	--	--	--
22	Mrg-0092	7-7-2004	7	309	--	45	--	288	--	--	9	--
23	Mrg-0149	11-4-2004	--	--	--	534	--	322	--	--	--	--
24	Mrg-0147	11-3-2004	--	--	--	216	--	55.8	--	--	--	--
25	Mrg-0112	8-10-2004	2	340	<0.4	6,070	<.08	1,510	.51	29.5	E3	450
26	Mrg-0138	9-23-2004	--	--	--	2,490	--	124	--	--	--	--
27	Mrg-0141	10-20-2004	--	--	--	610	--	925	--	--	--	--
28	Mrg-0176	5-6-2005	3	223	<0.4	462	<.08	230	1.76	132	10	--
29	Mrg-0159	4-15-2005	--	--	--	24,100	--	3,920	--	--	--	--
30	Mrg-0091	7-7-2004	--	--	--	<6	--	32.9	--	--	--	--
31	Mrg-0179	5-11-2005	--	--	--	<6	--	<6	--	--	--	--
32	Mrg-0148	11-4-2004	--	--	--	<6	--	22.2	--	--	--	--

Appendix 2d. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: µg/L, micrograms per liter; pCi/L, picocuries per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Arsenic, total (µg/L)	Barium, total (µg/L)	Cadmium, dissolved (µg/L)	Iron, dissolved (µg/L)	Lead, dissolved (µg/L)	Manganese, dissolved (µg/L)	Nickel, dissolved (µg/L)	Strontium, dissolved (µg/L)	Zinc, total (µg/L)	Radon-222, total (pCi/L)
33	Mrg-0120	8-25-2004	--	--	--	131	--	98.6	--	--	--	--
34	Mrg-0116	8-12-2004	3	220	<.04	1,080	<.08	556	1.34	137	<4	410
35	Mrg-0155	4-13-2005	3	291	<.04	393	<.08	287	1.01	79.8	<6	80
36	Mrg-0106	8-3-2004	--	--	--	<6	--	4.2	--	--	--	--
37	Mrg-0170	4-28-2005	--	--	--	<6	--	E.5	--	--	--	--
38	Mrg-0100	7-15-2004	--	--	--	<6	--	9	--	--	--	--
39	Mrg-0085	5-19-2004	<4	76	² <.5	² 152	² <2	² <17	² <1	--	--	340
40	Mrg-0121	8-25-2004	--	--	--	14	--	10	--	--	--	--
41	Mrg-0160	4-19-2005	<2	137	--	12	--	210	--	--	<6	--
42	Mrg-0123	8-26-2004	E2	53.5	E.02	<6	.26	E.7	3.4	1,140	8	--
43	Mrg-0113	8-10-2004	<2	37.5	<.04	62	.11	19.3	1.8	219	6	1,330
44	Mrg-0088	7-1-2004	--	--	--	<6	--	<.8	--	--	--	--
45	Mrg-0118	8-24-2004	2	52.9	<.04	7,050	<.08	118	30.6	51.3	191	10
46	Mrg-0110	8-5-2004	--	--	--	37	--	8.7	--	--	--	350
47	Mrg-0167	4-25-2005	<2	48.8	--	E5	--	0.6	--	--	<6	--
48	Mrg-0097	7-13-2004	--	--	--	<6	--	<.8	--	--	--	--
49	Mrg-0117	6-29-2004	--	--	--	44	--	2.4	--	--	--	--
50	Mrg-0133	9-9-2004	--	--	--	831	--	274	--	--	--	--
51	Mrg-0171	5-2-2005	--	--	--	<6	--	2.1	--	--	--	--
52	Mrg-0172	5-2-2005	--	--	--	18	--	103	--	--	--	--
53	Mrg-0173	5-3-2005	<2	16.9	<.04	129	0.15	118	3.21	588	<6	100
54	Mrg-0103	7-28-2004	--	--	--	263	--	1,300	--	--	--	--
55	Mrg-0102	7-28-2004	--	--	--	4,900	--	1,060	--	--	--	--
56	Mrg-0101	7-27-2004	--	--	--	10	--	240	--	--	--	--
57	Mrg-0105	8-3-2004	--	--	--	19	--	289	--	--	--	--
58	Mrg-0093	7-8-2004	3	227	--	517	--	636	--	--	5	--
59	Mrg-0107	8-4-2004	--	--	--	89	--	6.7	--	--	--	--
60	Mrg-0131	9-8-2004	E1	128	<.04	<6	.29	12.2	1.31	92.9	E4	170
61	Mrg-0132	9-8-2004	--	74.7	<.04	<6	E.05	1	1.67	5,520	E3	490
62	Mrg-0128	9-1-2004	--	--	--	10	--	16.1	--	--	--	--
63	Mrg-0125	8-31-2004	--	--	--	28	--	60.9	--	--	--	--
64	Mrg-0126	8-31-2004	--	--	--	68	--	102	--	--	--	--
65	Mrg-0127	9-1-2004	--	--	--	1,590	--	346	--	--	--	--

Appendix 2d. Water-quality data values for wells, Morgan County, West Virginia. —Continued

[USGS, U.S. Geological Survey. Units of measure: µg/L, micrograms per liter, pCi/L, picocuries per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Arsenic, total (µg/L)	Barium, total (µg/L)	Cadmium, dissolved (µg/L)	Iron, dissolved (µg/L)	Lead, dissolved (µg/L)	Manganese, dissolved (µg/L)	Nickel, dissolved (µg/L)	Strontium, dissolved (µg/L)	Zinc, total (µg/L)	Radon-222, total (pCi/L)
66	Mrg-0096	7-13-2004	--	--	--	<6	--	13.9	--	--	--	--
67	Mrg-0108	8-4-2004	--	--	--	13	--	129	--	--	--	--
68	Mrg-0129	9-2-2004	--	--	--	42	--	125	--	--	--	560
69	Mrg-0109	8-4-2004	--	--	--	E5	--	10.2	--	--	--	--
70	Mrg-0119	8-24-2004	4	328	<.04	126	<.08	123	.8	1,290	10	380
71	Mrg-0099	7-14-2004	--	--	--	<6	--	1.1	--	--	--	--
72	Mrg-0098	7-14-2004	--	--	--	<6	--	<.8	--	--	--	--
73	Mrg-0168	4-26-2005	<2	98.8	--	2,760	--	110	--	--	E6	--
74	Mrg-0111	8-9-2004	--	--	--	77	--	120	--	--	--	--
75	Mrg-0169	4-26-2005	2	49.1	--	1,120	--	1,050	--	--	23	--
76	Mrg-0158	4-14-2005	--	--	--	1,380	--	534	--	--	--	--
77	Mrg-0089	7-1-2004	--	--	--	42	--	135	--	--	--	--
78	Mrg-0087	6-30-2004	--	--	--	<6	--	179	--	--	--	--
79	Mrg-0166	4-22-2005	E1	21.2	--	E5	--	8.6	--	--	13	--
80	Mrg-0175	5-5-2005	8	390	<.04	1,140	<.08	382	1.96	36.1	12	--
81	Mrg-0086	5-19-2004	8	150	² <.5	³ 3,370	--	² 1,190	² <1	--	3	250
82	Mrg-0154	4-12-2005	<2	25.9	<.04	<6	.95	E.5	2.14	9.69	8	--
83	Mrg-0164	4-21-2005	--	--	--	404	--	358	--	--	--	--
84	Mrg-0174	5-3-2005	<2	101	.11	8	E.05	14	3.32	28.5	696	640
85	Mrg-0163	4-20-2005	<2	47.8	--	8	--	85.1	--	--	8	70
86	Mrg-0165	4-21-2005	E1	26.3	--	E3	--	4.7	--	--	31	--
87	Mrg-0134	9-10-2004	--	--	--	855	--	213	--	--	--	--
88	Mrg-0115	8-11-2004	--	--	--	60	--	239	--	--	--	--
89	Mrg-0153	4-12-2005	--	--	--	E6	--	11.1	--	--	--	--
90	Mrg-0162	4-20-2005	<2	38	--	<6	--	2.7	--	--	E3	--
91	Mrg-0150	3-17-2005	--	--	--	89	--	220	--	--	--	--

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0090 is the 90th well site established in Morgan County, West Virginia.

² Sample concentration reported as "Total."

Appendix 3a. Water-quality data values for springs, Morgan County, West Virginia.

[USGS, U.S. Geological Survey. Units of measure: mg/L, milligrams per liter; µg/L, micrograms per liter; col/100 ml, colonies per 100 milliliters. Symbols: --, not determined; <, less than; E, estimated]

Site number	USGS station name	Date	Nitrate+ nitrite, dissolved (mg/L as N)	Nitrite, dissolved (mg/L as N)	Ortho-phosphate, dissolved (mg/L)	Ortho-phosphate, dissolved (mg/L as P)	Phosphorus, dissolved (mg/L)	E. coli, (col/100 ml)	Coliform, fecal (col/100 ml)	Coliform, total (col/100 ml)	Iron, dissolved (µg/L)	Manganese, dissolved (µg/L)
93	Mrg-0058	5/10/2005	<0.06	<0.008	--	<0.02	<0.04	<1	--	E3	2,010	212
94	Mrg-0012S	3/15/2005	.11	<.008	--	<.02	<.04	<1	<1	<1	<6	<.6
95	Mrg-0060S	5/9/2005	.31	<.008	--	<.02	<.04	E4	--	67	<6	112
96	Mrg-0177S	5/10/2005	.39	<.008	--	<.02	<.04	<1	--	21	<6	E.6
97	Mrg-0151S	3/16/2005	.21	<.008	--	<.02	<.04	--	<1	--	<6	<.6
97	Mrg-0151S	5/9/2005	--	--	--	--	--	<1	--	E9	--	--
98	Mrg-0017S	3/16/2005	.4	<.008	--	<.02	<.04	--	E1	--	<6	<.6
98	Mrg-0017S	5/9/2005	--	--	--	--	--	E2	--	E11	--	--
99	Mrg-0075S	3/15/2005	.66	<.008	--	<.02	<.04	<1	<1	E8	E4	5.6
100	Mrg-0042	3/15/2005	.26	<.008	0.077	.03	E.04	<1	<1	<1	<6	<.6

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0012S is the 12th spring site established in Morgan County, West Virginia.

Appendix 3b. Water-quality data values for springs, Morgan County, West Virginia.

[USGS, U.S. Geological Survey. Units of measure: mg/L, milligrams per liter. Symbols: --, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Magnesium, dissolved (mg/L)	Potassium, dissolved (mg/L)	Sodium, dissolved (mg/L)	Chloride, dissolved (mg/L)	Fluoride, dissolved (mg/L)	Silica, dissolved (mg/L)	Sulfate, dissolved (mg/L)	Solids, residue at 180 deg C dissolved (mg/L)	Ammonia, dissolved (mg/L as NH4)	Ammonia, dissolved (mg/L as N)
93	Mrg-0058	5/10/2005	7.21	0.62	6.53	3.84	0.1	21.1	33.8	165	0.09	0.07
94	Mrg-0012S	3/15/2005	7.15	.67	1.49	1.05	.1	9.3	5.8	189	--	<.04
95	Mrg-0060S	5/9/2005	9.1	.65	8	1.02	E.1	5.4	11.4	89	--	<.04
96	Mrg-0177S	5/10/2005	20.4	.67	3.06	.99	<.1	6.4	8.1	358	--	<.04
97	Mrg-0151S	3/16/2005	9.3	.57	2.62	6.03	.1	7.8	3.5	197	--	<.04
97	Mrg-0151S	5/9/2005	--	--	--	--	--	--	--	--	--	--
98	Mrg-0017S	3/16/2005	6.71	.96	2.3	4.43	E.1	6.7	6.7	137	--	<.04
98	Mrg-0017S	5/9/2005	--	--	--	--	--	--	--	--	--	--
99	Mrg-0075S	3/15/2005	2.38	.9	.83	1.36	<.1	5.7	11.2	31	--	<.04
100	Mrg-0042	3/15/2005	4.67	.95	4.64	2.37	E.1	9.9	14.6	161	--	<.04

¹USGS station same designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0012S is the 12th spring site established in Morgan County, West Virginia.

Appendix 3c. Water-quality data values for springs, Morgan County, West Virginia.

[USGS, U.S. Geological Survey. Units of measure: mg/L, milligrams per liter; Deg C, degrees Celsius; mv, millivolts; $\mu\text{S}/\text{cm}$, microsiemens per centimeter at 25°C ; NTU, nephelometric turbidity units. Symbols: --, not determined; <, less than; E, estimated]

Site number	¹ USGS station name	Date	Land-surface altitude (feet above mean sea level)	Redox potential (mv)	Turbidity (NTU)	Oxygen, dissolved (mg/L)	pH (units)	Specific conductance ($\mu\text{S}/\text{cm}$)	Temperature, (Deg C)	Water hardness (mg/L as CaCO_3)	Calcium, dissolved (mg/L)
93	Mrg-0058	5/10/2005	925	196	<1.0	7.8	6.9	258	9.9	130	40.8
94	Mrg-0012S	3/15/2005	900	140	<1.0	3.2	7.2	416	11.2	190	63
95	Mrg-0060S	5/9/2005	560	260	3.6	10.1	7.3	218	10	64	10.5
96	Mrg-0177S	5/10/2005	872	217	<1.0	4.3	6.7	814	10.8	390	124
97	Mrg-0151S	3/16/2005	820	148	<1.0	3.1	7.6	441	11.8	200	62.9
97	Mrg-0151S	5/9/2005	820	221	1.1	2.9	7	450	11.9	--	--
98	Mrg-0017S	3/16/2005	820	152	<1.0	6	7.4	323	9.9	130	39.7
98	Mrg-0017S	5/9/2005	820	232	<1.0	5.9	6.8	332	11	--	--
99	Mrg-0075S	3/15/2005	1,000	184	<1.0	6	5.4	59	10.5	18	3.33
100	Mrg-0042	3/15/2005	620	125	<1.0	4.3	7.1	366	21.7	140	48.8

¹USGS station name designates the chronological order in which sampling sites were established in a given county; for example, Mrg-0012S is the 12th spring site established in Morgan County, West Virginia.

Carol J. Boughton and Kurt J. McCoy—**Hydrogeology, Aquifer Geochemistry, and Ground-Water Quality in Morgan County, West Virginia—**
Scientific Investigations Report 2006-5198