

In cooperation with the San Antonio Water System

Geologic, Hydrologic, and Geochemical Identification of Flow Paths in the Edwards Aquifer, Northeastern Bexar and Southern Comal Counties, Texas



Scientific Investigations Report 2007–5285

Cover: U.S. Geological Survey hydrologic technicians preparing to collect samples on Landa Lake, New Braunfels, Texas.

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By Cassi L. Otero

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**U.S. Department of the Interior
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Datums and Abbreviations

Vertical coordinate information is referenced to the National Geodetic Vertical Datum of 1929 (NGVD 29) or the North American Vertical Datum of 1988 (NAVD 88).

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above the vertical datum.

Specific conductance is given in microsiemens per centimeter at 25 degrees Celsius ($\mu\text{S}/\text{cm}$).

Concentrations of chemical constituents in water are given in either milligrams per liter (mg/L) or micrograms per liter ($\mu\text{g}/\text{L}$).

Geologic, Hydrologic, and Geochemical Identification of Flow Paths in the Edwards Aquifer, Northeastern Bexar and Southern Comal Counties, Texas

By Cassi L. Otero

Abstract

The U.S. Geological Survey, in cooperation with the San Antonio Water System, conducted a 4-year study during 2002–06 to identify major flow paths in the Edwards aquifer in northeastern Bexar and southern Comal Counties (study area). In the study area, faulting directs ground water into three hypothesized flow paths that move water, generally, from the southwest to the northeast. These flow paths are identified as the southern Comal flow path, the central Comal flow path, and the northern Comal flow path. Statistical correlations between water levels for six observation wells and between the water levels and discharges from Comal Springs and Hueco Springs yielded evidence for the hypothesized flow paths. Strong linear correlations were evident between the datasets from wells and springs within the same flow path and the datasets from wells in areas where flow between flow paths was suspected. Geochemical data (major ions, stable isotopes, sulfur hexafluoride, and tritium and helium) were used in graphical analyses to obtain evidence of the flow path from which wells or springs derive water. Major-ion geochemistry in samples from selected wells and springs showed relatively little variation. Samples from the southern Comal flow path were characterized by relatively high sulfate and chloride concentrations, possibly indicating that the water in the flow path was mixing with small amounts of saline water from the freshwater/saline-water transition zone. Samples from the central Comal flow path yielded the most varied major-ion geochemistry of the three hypothesized flow paths. Central Comal flow path samples were characterized, in general, by high calcium concentrations and low magnesium concentrations. Samples from the northern Comal flow path were characterized by relatively low sulfate and chloride concentrations and high magnesium concentrations. The high magnesium concentrations characteristic of northern Comal flow path samples from the recharge zone in Comal County might indicate that water from the Trinity aquifer is entering the Edwards aquifer in the subsurface. A graph of the relation between the stable isotopes deuterium and delta-18

oxygen showed that, except for samples collected following an unusually intense rain storm, there was not much variation in stable isotope values among the flow paths. In the study area deuterium ranged from -36.00 to -20.89 per mil and delta-18 oxygen ranged from -6.03 to -3.70 per mil. Excluding samples collected following the intense rain storm, the deuterium range in the study area was -33.00 to -20.89 per mil and the delta-18 oxygen range was -4.60 to -3.70 per mil. Two ground-water age-dating techniques, sulfur hexafluoride concentrations and tritium/helium-3 isotope ratios, were used to compute apparent ages (time since recharge occurred) of water samples collected in the study area. In general, the apparent ages computed by the two methods do not seem to indicate direction of flow. Apparent ages computed for water samples in northeastern Bexar and southern Comal Counties do not vary greatly except for some very young water in the recharge zone in central Comal County.

Introduction

The Edwards aquifer is the main source of public water supply for the city of San Antonio, Texas, and the surrounding area and provides nearly all of the water for industrial, military, and irrigation use in the region (fig.1). Withdrawals from the aquifer to meet San Antonio's increasing water-supply needs might be a threat to minimum mandated sustained flows at Comal Springs (Edwards Aquifer Authority, 2007), the largest spring in the Southwest. The springs supply water to downstream users, sustain federally-listed endangered species, and support local economies through tourism. Increased knowledge of the complex hydrologic processes that control water availability in the Edwards aquifer in the vicinity of Comal Springs is imperative for optimal resource management. The U.S. Geological Survey (USGS), in cooperation with the San Antonio Water System (SAWS), conducted a 4-year study during 2002–06 in northeastern Bexar County and southern Comal County to identify flow paths in the Edwards aquifer. The study area (fig. 1) includes small parts

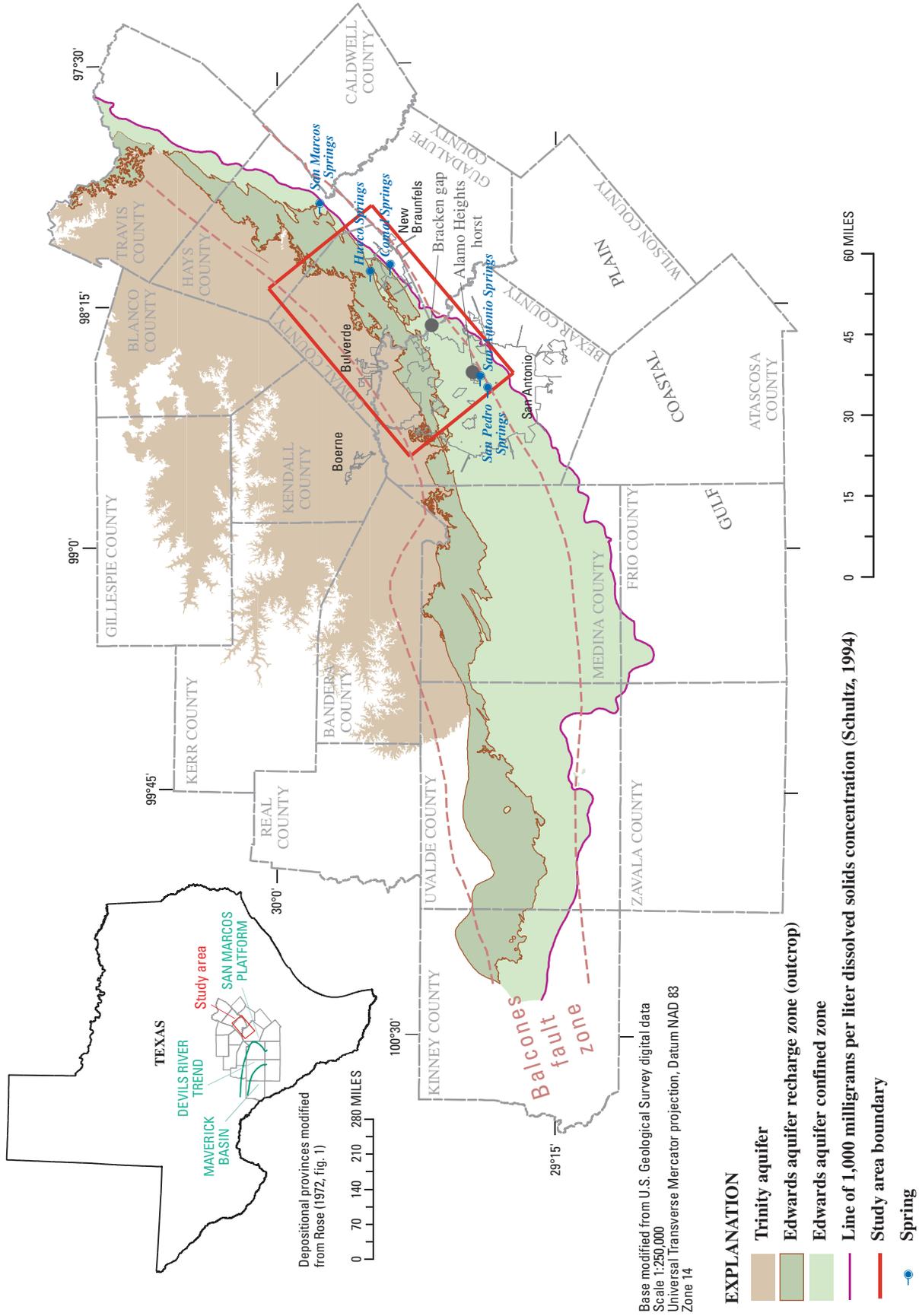


Figure 1. San Antonio segment of the Edwards aquifer, location of study area, and relation between Edwards and Trinity aquifers, south-central Texas.

of Guadalupe and Hays Counties in addition to northeastern Bexar County and southern Comal County.

Purpose and Scope

The purpose of this report is to describe major ground-water flow paths in the Edwards aquifer in northeastern Bexar and southern Comal Counties identified on the basis of geologic, hydrologic, and geochemical data. Hydrogeology and geologic structure based on work done by Maclay and Small (1984), Small (1986), Small and Hanson (1994), Maclay (1995), and Stein and Ozuna (1994) and flow path work done by Maclay and Land (1988) and Groschen (1996) provided a basis for the initial selection of hypothesized flow paths in the study area. Flow paths were further defined using hydrologic data collected from water-level observation wells and springflow monitoring sites in the study area. Altitudes of the potentiometric surface within each flow path provided evidence of flow directions. Comparisons of water-level and spring-discharge hydrographs and statistical correlation of water levels and spring discharges were used to show relations between water levels at wells and spring discharges within flow paths. Continuous water-level data were collected at six observation wells from mid-March 2004 through September 2006. Discharge data were obtained from two springflow sites for the same time period. Ground-water chemistry and isotope data were compiled from samples collected from 76 wells and nine springs (and spring orifices of major springs) during 1996–2006. Isotope data also were obtained from samples collected at four rainfall sites during 1998–2003.

Description of the Edwards Aquifer and Study Area

The San Antonio segment of the Edwards aquifer (hereinafter, Edwards aquifer) comprises Lower Cretaceous-age rocks of the Edwards Group (Rose, 1972) and the Georgetown Formation. The Edwards Group in the study area comprises two stratigraphic units, the basal Kainer Formation and the upper Person Formation. Each of those units comprises several informal members. The basal member of the Person Formation is a laterally extensive marine deposit consisting of poorly permeable, dense, carbonate mudstone known as the regional dense member (Maclay, 1995).

Most recharge to the Edwards aquifer occurs in the recharge zone (aquifer outcrop) west of Bexar County (fig. 1), where streams originating north of the aquifer flow across and lose most or all of their flows into highly faulted and fractured limestone. Additional recharge enters the aquifer through the recharge zone in Bexar, Comal, and Hays Counties. After the water enters the aquifer, it moves generally in an easterly direction to discharge points in Bexar County, mainly municipal water-supply wells. Water not discharged to wells then continues generally toward the northeast along and parallel to northeast-trending faults in the study area to discharge

points in Comal and Hays Counties, primarily Comal Springs in Comal County and San Marcos Springs in Hays County (fig. 1).

The study area is in an extensively faulted section of Cretaceous strata known as the Balcones fault zone (fig. 1). The fault zone developed as a result of extensional faulting and is characterized by a network of en-echelon, high-angle, mostly down-to-the-coast normal faults along the northwestern margin of the Gulf Coastal Plain (Maclay and Small, 1984; Maclay, 1995). The Cretaceous strata were vertically displaced, intensively fractured, and differentially rotated within a series of southwest-to-northeast trending fault blocks (Barker and Ardis, 1996). The fault blocks, and their subsequent erosion and dissolution, are major factors affecting flow in the aquifer.

Maclay and Land (1988, fig. 22) defined four major flow units in the Edwards aquifer. The flow units originate in areas referred to as storage units in the recharge zone and are regions of confined flow that generally move water initially to the southwest and then to the east and northeast to discharge at major springs. The eastern flow unit described by Maclay and Land (1988) originates in the study area. The western-southern, south-central, and north-central flow units coalesce in northeastern Bexar County and southern Comal County in the vicinity of Cibolo Creek and Interstate Highway 35.

Maclay and Small (1984, p. 50) estimated transmissivities for the Edwards aquifer to range from 200,000 to 2,000,000 feet squared per day. Maclay and Small (1984) describe one of the most transmissive areas in the Edwards aquifer as occurring within a narrow, northeast-trending band downgradient from the area of coalescence of the three southernmost flow units. This band of high transmissivity (fig. 2) (labeled “R” in Maclay and Small [1984, fig. 20]) is bounded on the northwest and southeast by faults where less-permeable rocks of the upper confining unit of the aquifer are juxtaposed against rocks of the Edwards aquifer.

Recharge to and flow within the Edwards aquifer in northeastern Bexar and southern Comal Counties are complicated by the structure and stratigraphy of the rocks. The study area (fig. 1) is on the structural high known as the San Marcos Platform (Rose, 1972). The San Marcos Platform is extensively faulted in the study area. The thickness of the Edwards aquifer in the study area is about 450 feet (Small and Hanson, 1994, p. 5; Stein and Ozuna, 1994, p. 5). Three major faults, the Comal Springs fault and the Hueco Springs fault in southern Comal County and the Northern Bexar fault (Maclay and Land, 1988) in northeastern Bexar County, are potentially effective barriers to flow in the Edwards aquifer in the study area (fig. 2).

The Comal Springs complex (fig. 3) issues from the Comal Springs fault, which has as much as 500 feet of offset (Maclay and Land, 1988, p. A42) and is juxtaposed against the younger and less permeable upper confining unit. The springs developed because a roughly north-south trending transverse fault east of New Braunfels completely offsets the Edwards aquifer in the downthrown block of the Comal Springs fault,

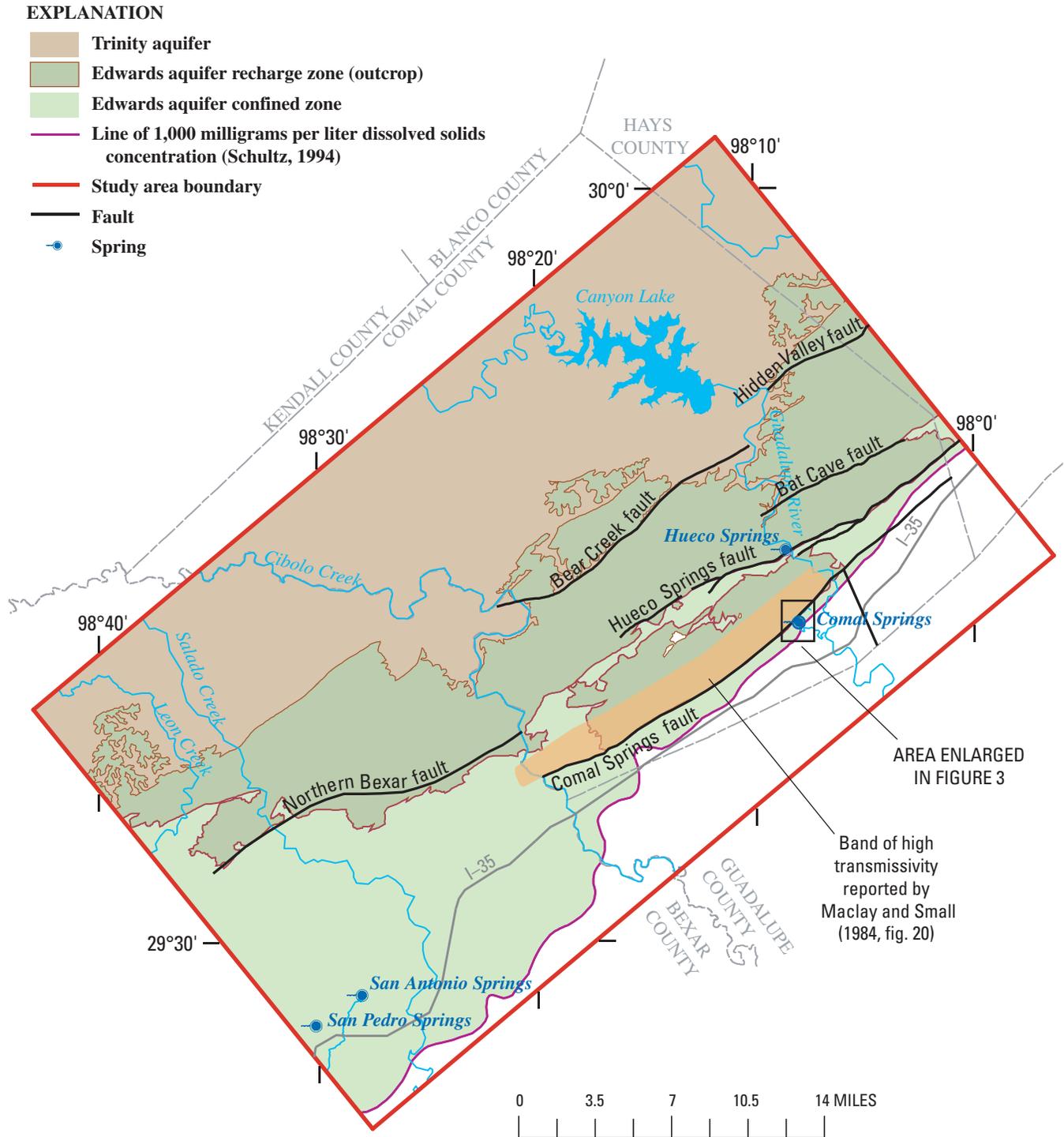


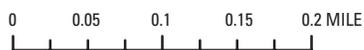
Figure 2. Study area and relation between recharge zone and confined zone of the Edwards aquifer and the Trinity aquifer, northeastern Bexar and southern Comal Counties, Texas.

creating a barrier to northeastward flow and forcing water upward along the Comal Springs fault (Klemm and others, 1979, fig. 8). About one-fourth of the springflow from the Comal Springs complex discharges from three large spring orifices (Comal 1, Comal 2, and Comal 3 on the west side of the complex) that are sourced in the upthrown block of the

Comal Springs fault (Ogden and others, 1985; LBG-Guyton Associates, 2004). The remaining springflow is discharged from numerous spring orifices and seeps that are within and near the banks of Landa Lake; these springs and seeps are sourced in the downthrown block of the Comal Springs fault.



Base from StreetMap USA, 2004



EXPLANATION

- Edwards aquifer recharge zone (outcrop)
- Edwards aquifer confined zone
- Comal Springs complex
- Line of 1,000 milligrams per liter dissolved solids concentration (Schultz, 1994)
- Spring

Figure 3. Locations of major springs in the Comal Springs complex, southern Comal County, Texas.

Hueco Springs issues from the Hueco Springs fault (fig. 2), which has 380 to 400 feet of offset in the vicinity of the springs (William F. Guyton and Associates, 1979, p. 25) and exposes much of the less-permeable upper confining unit on the downthrown block. The series of fault blocks immediately adjacent to and southeast of the Hueco Springs fault in southern Comal County are oriented mainly down-to-the-northwest (fig. 4). Consequently, this section of Edwards aquifer between the Hueco Springs fault and the Comal Springs fault is in a graben that is tilted opposite to the prevailing structure in the area. The fault blocks near Hueco Springs fault generally contain the entire thickness of the Edwards aquifer and parts of the upper confining unit, whereas the fault blocks near Comal Springs fault generally contain incomplete, unconfined sections of the Edwards aquifer exposed at the surface. The fault blocks contain numerous transverse faults that possibly impede ground-water flow. Northwest of Hueco Springs fault are faults that probably are not effective barriers to flow, including Bat Cave, Bear Creek, and Hidden Valley faults (fig. 2).

Acknowledgments

Thanks are extended to SAWS and other well owners who granted permission to have their wells sampled for water quality and measured for water levels. Special thanks are extended to Hanson Aggregates, Inc., and New Braunfels Utility (NBU) for access to their wells for use as real-time, continuous water-level observation sites for this study.

Methods of Investigation

Data Collection and Sample Analysis

Six wells were used for the collection of continuous water-level data (fig. 5; table 1). Two real-time, continuous water-level data-collection sites—DX-68-22-913 (Hanson) and DX-68-23-502 (Solms)—were established to collect hourly ground-water-level data for this study (appendix 1.1, 1.2). Water-level data also were compiled on an hourly basis from four existing USGS real-time, continuous ground-water-level observation wells—AY-68-29-103 (HCV), AY-68-37-203 (Bexar), DX-68-23-304 (NBU-LCRA), and DX-68-30-208 (Bracken). Collection of continuous water-level data at the observation wells began in mid-March 2004 and continued through September 2005 at wells HCV and Bracken and through September 2006 at the four other wells. USGS-computed springflow data for Comal Springs and Hueco Springs also were compiled for mid-March 2004 through September 2006. All water-level and spring-flow data are in the USGS National Water Information System for Texas (U.S. Geological Survey, 2006a).

Ground-water-chemistry and isotope data were collected and compiled from 76 wells and nine springs (and spring orifices from major springs) during 1996–2006 (fig. 5; table 1). Data were collected for this study during 2003 (and during 2006 at selected sites). Additional data were collected during 1996–2006 as part of the National Water Quality Assessment (NAWQA) program (U.S. Geological Survey, 2006c). Two wells and two springs provided data from both sampling periods. Sixty-three wells and two springs provided NAWQA data only. Eleven wells and five springs provided data only for this study.

Samples were collected for dissolved gases (methane, carbon dioxide, oxygen, nitrogen, and argon), sulfur hexafluoride (SF_6), and tritium (^3H) and helium-3 (^3He) concentrations from 13 wells and seven springs during 2003 (table 1; appendix 1.3). $^3\text{H}/^3\text{He}$ ratios were computed and used to estimate an apparent age (year sampled minus recharge year) of the water. Four wells (DX-68-22-810, DX-68-22-905, DX-68-23-304, and DX-68-23-710) and six springs (Comal 1, Comal 3, Comal 5, Comal 7, Comal-Spring Island, and Hueco A) were resampled in 2006 for $^3\text{H}/^3\text{He}$ because samples collected in 2003 resulted in either inconclusive or questionable age dates for these wells and springs. The 2006 sample for well DX-68-22-810 resulted in an inconclusive age date.

Isotope data were collected and compiled from four rainfall sites (fig. 5, table 1). Two sites provided only data collected in 1998 as part of the NAWQA program (U.S. Geological Survey, 2006c). The remaining two sites provided both NAWQA data and 2003 data.

Water-chemistry data included field properties (water temperature, pH, specific conductance, and dissolved oxygen), major ions, trace elements, nutrients, and alkalinity. Water samples were collected, processed, and preserved using standard USGS protocols as described in Wilde and others (1999, 2003, and 2004). The concentrations of major ions, trace elements, and nutrients in the water samples were measured by the USGS National Water Quality Laboratory in Denver, Colo., using approved methods (Fishman and Friedman, 1989; Patton and Truitt, 1992, 2000; Faires, 1993; Fishman, 1993; American Public Health Association, 1998; Garbarino and others, 2006).

The ratios of naturally occurring, stable isotopes of hydrogen (^2H) and oxygen ($^{18/16}\text{O}$) were measured by the USGS Stable Isotope Laboratory in Reston, Va., using approved methods (U.S. Geological Survey, 2005). Results for stable isotope analysis are reported as delta deuterium (δD) and delta 18-oxygen ($\delta^{18}\text{O}$), which represent the relative difference in parts per thousand (per mil) between the sample isotope ratio and the isotope ratio of a known standard (Kendall and McDonnell, 1998).

Dissolved gases and SF_6 samples were analyzed by the USGS Chlorofluorocarbon Laboratory in Reston, Va., using approved methods (U.S. Geological Survey, 2006b). $^3\text{H}/^3\text{He}$ samples collected in 2003 were analyzed by personnel at the Nobel Gas Laboratory of Lamont-Doherty Earth Observatory in Palisades, N.Y., and samples collected in 2006 were

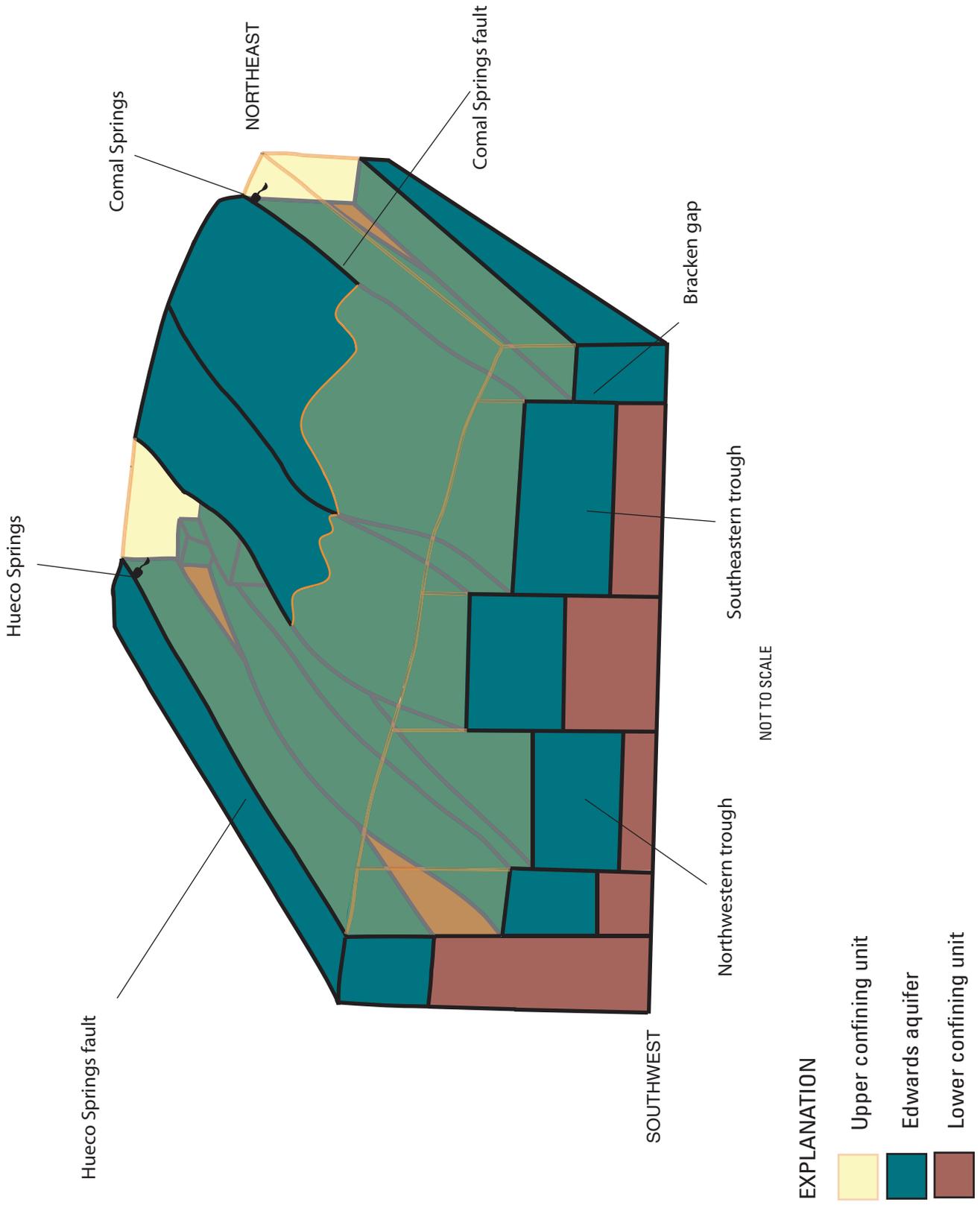


Figure 4. Idealized block diagram of the Edwards aquifer between Comal Springs fault and Hueco Springs fault, southern Comal County, Texas.

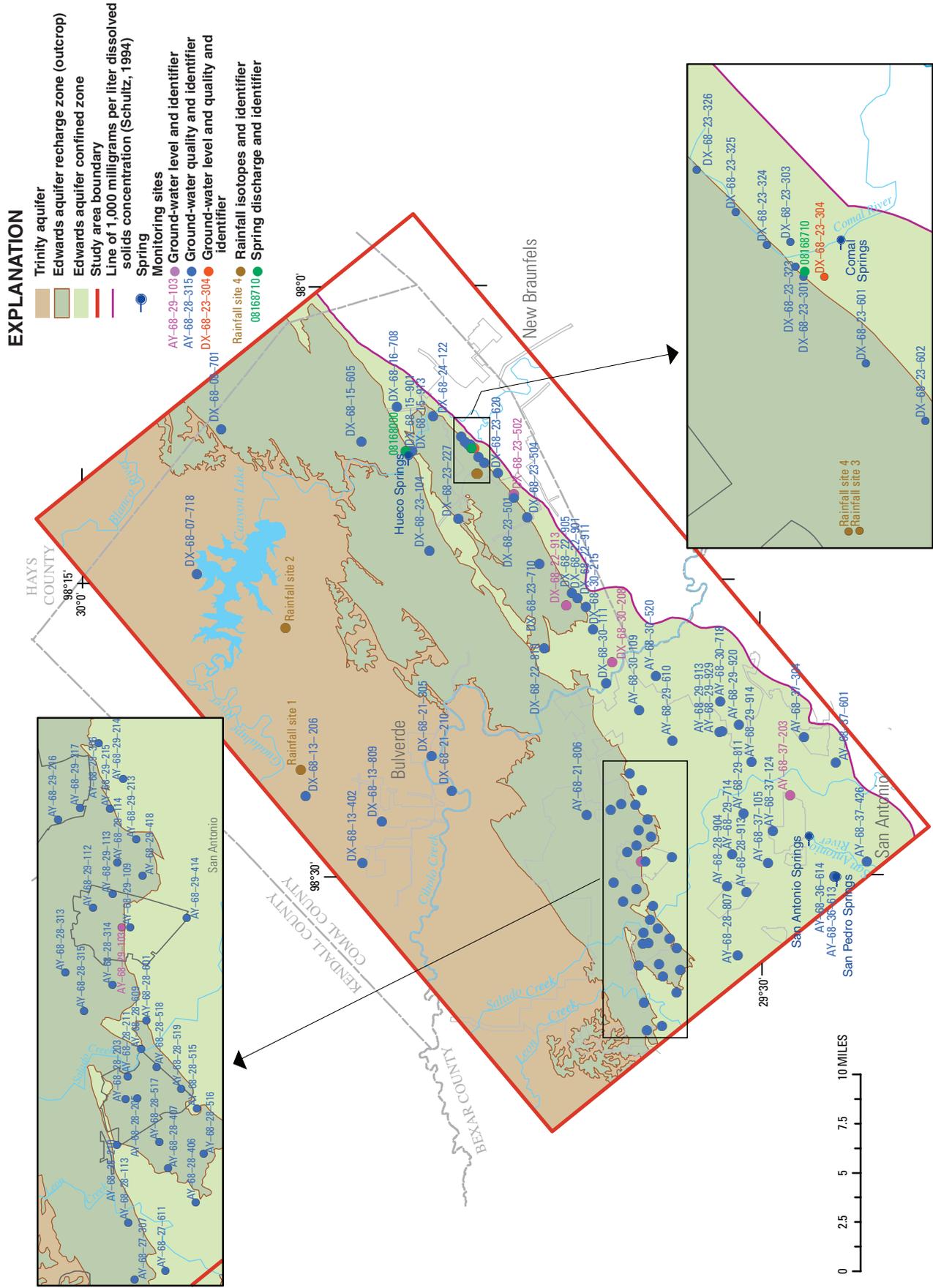


Figure 5. Locations of wells and spring orifices from which water-quality samples were collected and locations of continuous water-level observation wells, northeastern Bexar and southern Comal Counties, Texas.

Table 1. Ground-water sites (wells, springs, and spring orifices), springflow sites, and rainfall sites, northeastern Bexar and southern Comal Counties, Texas, from which data were collected and compiled for this report.

[Datum for open interval is land surface; USGS, U.S. Geological Survey; NA, not applicable; P, perforated or slotted; QW, water chemistry; I, deuterium and oxygen isotopes; X, open hole; WLC, continuous water level; DG, dissolved gas; SF, sulfur hexafluoride; THE, tritium and helium isotopes; W, walled; S, screen; --, unknown; R, wire-wound; DIS, discharge]

USGS site number	State well number (fig. 5)	USGS name	Aquifer	Total depth (feet)	Open interval (feet)	Completion type	Data type	Period of record
Ground-water sites (wells, springs, and spring orifices)								
293746098265401	AY-68-21-806	NA	Edwards	290	225-290	P	QW, I	1998-2006
293508098375101	AY-68-27-307	NA	Edwards	270	190-270	P	QW, I	1998-2006
293429098373801	AY-68-27-611	NA	Edwards	260	190-260	P	QW, I	1998-2006
293516098362801	AY-68-28-113	NA	Edwards	320	250-320	P	QW, I	1998-2006
293518098332601	AY-68-28-203	NA	Edwards	435	162-435	X	QW, I	2005-06
293504098332601	AY-68-28-205	NA	Edwards	485	273-485	X	QW, I	1996
293530098343401	AY-68-28-210	NA	Edwards	281	201-281	P	QW, I	1998-2006
293516098325501	AY-68-28-211	NA	Edwards	300	220-300	P	QW, I	1998-2006
293635098302301	AY-68-28-313	NA	Edwards	300	197-297	P	QW, I	1998-2006
293535098304101	AY-68-28-314	NA	Edwards	241	170-240	P	QW, I	1998-2006
293611098311901	AY-68-28-315	NA	Edwards	280	190-280	P	QW, I	1998-2006
293550098355801	AY-68-28-406	NA	Edwards	310	224-304	P	QW, I	1998-2006
293425098350801	AY-68-28-407	NA	Edwards	310	240-310	P	QW, I	1998-2006
293348098334101	AY-68-28-515	NA	Edwards	305	245-305	P	QW, I	1998-2006
293340098344701	AY-68-28-516	NA	Edwards	302	202-302	P	QW, I	1998-2006
293436098343001	AY-68-28-517	NA	Edwards	261	181-261	P	QW, I	1998-2006
293439098324101	AY-68-28-518	NA	Edwards	261	181-261	P	QW, I	1998-2006
293408098331301	AY-68-28-519	NA	Edwards	280	179-280	P	QW, I	1998-2006
293451098313201	AY-68-28-601	NA	Edwards	500	40-500	X	QW	2005
293459098321401	AY-68-28-609	NA	Edwards	260	200-260	P	QW, I	1998-2006
293111098340901	AY-68-28-807	NA	Edwards	685	420-685	X	QW	2005
293133098303201	AY-68-28-904	NA	Edwards	640	528-640	X	QW, I	1997
293042098305201	AY-68-28-913	NA	Edwards	787	448-787	X	QW	2005
293522098291201	AY-68-29-103	HCV	Edwards	547	90-547	X	WLC	2004-05
293512098291701	AY-68-29-109	NA	Edwards	600	197-590	X	QW, I	1996-2006
293559098284801	AY-68-29-112	NA	Edwards	260	160-260	P	QW, I	1998-2006
293534098282801	AY-68-29-113	NA	Edwards	241	160-240	P	QW, I	1998-2006
293528098274301	AY-68-29-114	NA	Edwards	201	120-201	P	QW, I	1998-2006
293504098270901	AY-68-29-213	NA	Edwards	222	142-222	P	QW, I	1998-2006
293520098254101	AY-68-29-214	NA	Edwards	222	101-221	P	QW, I	1998-2006
293537098262401	AY-68-29-215	NA	Edwards	180	120-180	P	QW, I	1998-2006
293643098264001	AY-68-29-216	NA	Edwards	261	180-260	P	QW, I	1998-2006
293615098262301	AY-68-29-217	NA	Edwards	257	188-257	P	QW, I	1998-2006

Table 1. Ground-water sites (wells, springs, and spring orifices), springflow sites, and rainfall sites, northeastern Bexar and southern Comal Counties, Texas, from which data were collected and compiled for this report—Continued.

USGS site number	State well number (fig. 5)	USGS name	Aquifer	Total depth (feet)	Open interval (feet)	Completion type	Data type	Period of record
Ground-water sites (wells, springs, and spring orifices)—Continued								
293551098244801	AY-68-29-306	NA	Edwards	295	56-295	X	QW, I	1996-2006
293359098290301	AY-68-29-414	NA	Edwards	710	380-660	X	QW	2004
293456098280201	AY-68-29-418	NA	Edwards	181	111-181	P	QW, I	1998-2006
293358098231101	AY-68-29-610	NA	Edwards	811	359-811	X	QW, I, DG, SF, THE	2003-04
293120098285801	AY-68-29-714	NA	Edwards	848	408-848	X	QW	2004
293055098264301	AY-68-29-811	NA	Edwards	612	209-612	X	QW, I, DG, SF, THE	2003
293149098224201	AY-68-29-913	NA	Edwards	784	360-784	X	QW, I, DG, SF, THE	2003
293026098241401	AY-68-29-914	NA	Edwards	845	423-845	X	QW, I, DG, SF, THE	2003
293100098225401	AY-68-29-920	NA	Edwards	655	493-655	X	QW, I	1997
293145098224201	AY-68-29-929	NA	Edwards	870	409-870	X	QW	2004
293525098213701	AY-68-30-109	NA	Edwards	710	314-710	X	QW, I	1997
293440098195201	AY-68-30-520	NA	Edwards	777	427-777	X	QW, I, DG, SF, THE	2003
293119098211201	AY-68-30-718	NA	Edwards	877	530-877	X	QW	2004
292648098303701	AY-68-36-613	NA ¹	Edwards	NA	NA	NA	QW, I	1997
292648098303401	AY-68-36-614	NA ¹	Edwards	NA	NA	NA	QW, I	1997
292944098292301	AY-68-37-105	NA	Edwards	1050	476-1,050	X	QW	2004
292931098274601	AY-68-37-124	NA	Edwards	557	390-557	X	QW	2005
292845098255401	AY-68-37-203	Bexar	Edwards	874	491-874	X	WLC	2004-06
292808098230101	AY-68-37-304	NA	Edwards	685	600-685	X	QW, I	1997
292522098291901	AY-68-37-426	NA	Edwards	1114	774-1,114	X	QW	2004
292643098241801	AY-68-37-601	NA	Edwards	1150	1012-1,150	X	QW	2004
295458098143001	DX-68-07-718	NA	Trinity	180	59-180	W	QW, I	1996
295352098071201	DX-68-08-701	NA	Edwards	240	239-240	S	QW, I	1996-2006
295013098255201	DX-68-13-206	NA	Trinity	252	123-252	X	QW, I	1996-2002
294743098291801	DX-68-13-402	NA	Trinity	500	40-500	X	QW, I	1996
294650098265801	DX-68-13-809	NA	Trinity	580	100-580	W	QW, I	1996
294739098075301	DX-68-15-605	NA	Edwards	375	220-375	W	QW, I	1996-2006
294533098082301	DX-68-15-901	Hueco A ¹	Edwards	NA	NA	NA	QW, I, DG, SF, THE	2003-06
294533098082401	DX-68-15-913	Hueco B ¹	Edwards	NA	NA	NA	QW, I, DG, SF, THE	2003
294604098060801	DX-68-16-708	NA	Edwards	400	200-400	W	QW	2005
294344098253801	DX-68-21-210	NA	Trinity	464	--	--	QW, I	1996
294437098235201	DX-68-21-305	NA	Trinity	500	160-500	X	QW, I	1996
293937098182501	DX-68-22-810	NA	Edwards	405	--	--	QW, I, DG, SF, THE	2003-06
293807098155301	DX-68-22-901	NA	Edwards	255	148-255	X	QW	2005
293823098153801	DX-68-22-905	NA	Edwards	204	--	--	QW, I, DG, SF, THE	2003-06

Table 1. Ground-water sites (wells, springs, and spring orifices), springflow sites, and rainfall sites, northeastern Bexar and southern Comal Counties, Texas, from which data were collected and compiled for this report—Continued.

USGS site number	State well number (fig. 5)	USGS name	Aquifer	Total depth (feet)	Open interval (feet)	Completion type	Data type	Period of record
Ground-water sites (wells, springs, and spring orifices)—Continued								
293745098162001	DX-68-22-911	NA	Edwards	450	400-450	X	QW, I	1996-2006
293838098161601	DX-68-22-913	Hanson	Edwards	248	188-248	X	WLC	2004-06
294440098132701	DX-68-23-104	NA	Edwards	450	--	--	QW, I, DG, SE, THE	2003
294323098115101	DX-68-23-227	NA	Edwards	--	--	--	QW, I	1996-2006
294300098080001	DX-68-23-301	Comal 1 ¹	Edwards	NA	NA	NA	QW, I, DG, SE, THE	1997-2006
294249098080301	DX-68-23-303	NA	Edwards	1045	586-1,045	X	QW, I	1997
294239098081401	DX-68-23-304	NBU-LCRA	Edwards	1061	406-1,061	X	WLC, QW, I, DG, SE, THE	1998-2006
294248098081201	DX-68-23-323	Comal 3 ¹	Edwards	NA	NA	NA	QW, I, DG, SE, THE	2003-06
294255098080501	DX-68-23-324	Comal 7 ¹	Edwards	NA	NA	NA	QW, I, DG, SE, THE	2003-06
294304098075501	DX-68-23-325	Comal-Spring Island ¹	Edwards	NA	NA	NA	QW, I, DG, SE, THE	2003-06
294314098074101	DX-68-23-326	Comal 5 ¹	Edwards	NA	NA	NA	QW, I, DG, SE, THE	2003-06
294054098104601	DX-68-23-501	NA	Edwards	210	25-210	X	QW, I, DG, SE, THE	2003
294054098103501	DX-68-23-502	Solms	Edwards	448	46-448	X	WLC	2004-06
294019098114701	DX-68-23-504	NA	Edwards	215	119-215	X	QW	2005
294225098080301	DX-68-23-601	NA	Edwards	365	92-365	X	QW	2004
294206098090101	DX-68-23-602	NA	Edwards	790	149-239	P		
					239-790	X	QW, I, DG, SE, THE	2003
294137098093201	DX-68-23-620	NA	Edwards	--	--	--	QW, I	1996-2006
293949098141001	DX-68-23-710	NA	Edwards	217	--	--	QW, I, DG, SE, THE	2003-06
294428098063701	DX-68-24-122	NA	Edwards	--	--	--	QW, I	1996-2006
293653098201501	DX-68-30-111	NA	Edwards	230	--	--	QW, I, DG, SE, THE	2003
293636098190901	DX-68-30-208	Bracken	Edwards	292	220-240	R	WLC	2004-05
					272-292	P		
293729098173101	DX-68-30-215	NA	Edwards	660	185-660	X	QW	2005
Springflow sites								
08168710	NA	Comal Springs at New Braunfels, Tex. ¹	NA	NA	NA	NA	DIS	2004-06
08168000	NA	Hueco Springs near New Braunfels, Tex. ¹	NA	NA	NA	NA	DIS	2004-06
Rainfall sites								
295025098243200	NA	Rainfall site 1	NA	NA	NA	NA	I	1998
295104098171900	NA	Rainfall site 2	NA	NA	NA	NA	I	1998
294230098093300	NA	Rainfall site 3	NA	NA	NA	NA	I	1998-2003
294233098093300	NA	Rainfall site 4	NA	NA	NA	NA	I	1998-2003

¹ Spring or spring orifice.

analyzed by USGS personnel at the USGS Noble Gas Laboratory in Denver, Colo., using approved methods (U.S. Geological Survey, 2006b).

Quality Control and Quality Assurance of Geochemical Samples

Duplicate samples for major ions and trace elements were collected from one of the 13 wells (DX-68-23-710) sampled for this study (appendix 1.3). Duplicate samples for SF₆ were collected from all 13 wells sampled for this study. Duplicate samples for major ions, trace elements, and nutrients collected during 1996–2006 as part of the NAWQA program (U.S. Geological Survey, 2006c) were available for seven of the 65 wells supplying existing data for this study. Duplicate samples for major ions, trace elements, and nutrients were collected at one of the seven springs (DX-68-23-325) sampled for this study (appendix 1.3). Duplicate samples for SF₆ were collected from all seven springs sampled for this study.

Duplicate samples were used to evaluate the methods used by field and laboratory personnel to collect and analyze a given sample with consistent results. The non-zero relative percent differences between environmental and duplicate samples collected for major ions, trace elements, and nutrients for this study ranged from 0.41 percent for nickel in the spring sample to 54.5 percent for molybdenum in the spring sample with a median value of 3.1 percent. The greatest relative percent difference between the environmental and duplicate samples collected for SF₆ for this study was 10.3 percent in well DX-68-23-304.

Methods of Data Analysis

Hydrologic Data

Water levels measured in wells in the Edwards aquifer during October 30–November 3, 2000, in a multi-agency effort headed by the Edwards Aquifer Authority (Hamilton and Schindel, 2006) were used to construct a potentiometric-surface map for the study area. Water-level contours within each flow path were examined to determine flow directions in the Edwards aquifer.

Hydrographs of water levels at the six observation wells and discharge at Comal Springs and Hueco Springs were compared. Similarities in the water-level hydrographs might indicate either that the wells are responding to a wide-spread recharge event or that they are in the same flow path. Similarities in hydrographs for the spring discharge and water levels in wells could indicate that the spring and wells share the same flow path.

Statistical correlations between datasets of water levels for the six observation wells and discharge for Comal

Springs and Hueco Springs were analyzed using Pearson's *r* (Helsel and Hirsch, 1995). Pearson's *r* (or linear correlation coefficient) is a measure of the linear association between two variables. Pearson's *r* was calculated for each of the 28 combinations of datasets to indicate the strength of the linear association (correlations) between datasets. The assumption was that a strong linear correlation indicates a higher probability of a shared or common flow path than a nonlinear correlation or no correlation.

Hydrograph recession-curve analysis was done on spring-discharge hydrographs to identify the number and type of flow regimes (diffuse, fracture, and conduit) that characterize the ground-water flow path contributing to the spring. Hydrographs of spring discharge at Comal Springs and Hueco Springs were graphed, and recession curves (sections of the hydrograph where the discharge is falling after a sudden rise) spanning the study period were examined. Methods in Milanovich (1981), Bonacci (1993), Padilla and others (1994), Shevenell (1996), and Baedke and Krothe (2001) were used to examine the recession curves for breaks in the recession slope, which are indicative of a change from one flow regime to another within the karst continuum. A recession coefficient (α) was calculated for each part of the karst continuum. The value of α relates to the rate of release of water from the aquifer. In general, higher values of α indicate a steeper slope in the recession curve, and therefore, a release of water from conduit-type features in the aquifer, whereas lower values of α indicate a gentler recession slope and release of water from diffuse (matrix) features of the aquifer.

Geochemical Data

Geochemical data collected and compiled from selected wells and springs in the Edwards and Trinity aquifers in the study area were used to assess major-ion chemistry and apparent ground-water age (table 1). Ground-water flow paths were analyzed using geochemical and isotopic data. Additional ¹⁸O and ²H isotopic concentrations were estimated from rainfall samples to determine the local meteoric water line.

Dissolved gases and SF₆ were used to determine apparent age (year sampled minus recharge year) of ground-water samples. Apparent age is determined by comparing the concentration of SF₆ in the water sample to an established annual atmospheric concentration (Busenberg and Plummer, 2000). SF₆ is a trace gas in the atmosphere that accumulates in rainfall that eventually becomes recharge to the ground-water system. Mainly an anthropogenic compound, SF₆ also can occur naturally in fluid inclusions in some minerals and igneous rocks and in some volcanic and igneous fluids. Apparent age derived using this method does not take into account the mixing of young and old waters. The addition of excess air into the ground-water system, which can occur when air bubbles are dissolved during a rapid rise of the water table, increases the SF₆ concentration in the ground water to levels greater than the air-water equilibrium concentration. If the existence of

excess air is not considered in the computation of the recharge year, then the apparent age will be too young.

Apparent ages also were obtained using a method that involves measuring the relative abundance of tritium (^3H) and helium-3 (^3He) in a water sample. Tritium is a radioactive isotope of hydrogen that decays to ^3He with a half-life of 12.43 years. Tritium was introduced into the ground-water system in a series of peaks beginning in 1952, caused by atmospheric testing of nuclear devices, and reached a maximum concentration during 1963–64 (Schlosser and others, 1988). Although tritium concentrations in rainfall generally have decreased since this mid-1960s peak, an annual atmospheric concentration remains. Using a helium isotope mass balance, the amount of ^3He from the decay of ^3H (tritogenic ^3He denoted as $^3\text{He}^*$) is measured with the remaining ^3H in the water sample (Plummer and others, 2003). The values are added to determine the amount of ^3H that was present in the water sample at the time of recharge to the ground-water system. This calculated recharge concentration of ^3H then is compared to the established annual atmospheric concentration to determine an apparent age or recharge year for the water sample. This method of age-dating a water sample has proved to be effective for waters recharged within about the past 30 years and takes into account the mixing of young and old waters. Problems arise with this method when large amounts of terrigenous helium (derived from crustal or mantle sources) are present in a sample, such as in aquifers with host rock containing uranium or thallium, or in samples where young water has mixed with relatively old water containing terrigenous helium sources. In these cases, the ratio of $^3\text{He}/^4\text{He}$ for the terrigenous helium must be known within about 1 percent to determine an apparent age. If this ratio cannot be defined within the 1-percent limit, a range in age can be evaluated for a range in the terrigenous helium ratio (Schlosser and others, 1988).

Major-ion and stable isotope data were graphically evaluated to determine relations among constituents that could distinguish differences between flow paths in the study area. Piper trilinear diagrams were used to visually categorize the principal water compositions for each flow path (Hem, 1992). Scatterplots of constituents and ratios of constituents were constructed to graphically indicate variations in water samples collected from wells from different flow paths.

Geologic and Hydrologic Identification of Flow Paths

Southern Comal Flow Path

The southern Comal flow path (SCFP) (fig. 6) is bounded on the northwest by the Comal Springs fault and on the southeast by the freshwater/saline-water interface (threshold of 1,000-milligrams per liter (mg/L) dissolved solids concentration [Schultz, 1994]) in the Edwards aquifer (fig. 1). Although

the interface is not an actual physical barrier to flow, previous studies (Maclay and Small, 1984; Groschen, 1994; Lindgren and others, 2004) have indicated that flow in the transition zone (zone in which dissolved solids concentration ranges from 1,000 to 10,000 mg/L) is considerably more sluggish than flow in the immediately adjacent freshwater zone, likely because of relatively low permeability and transmissivity in the transition zone. The 1,000-mg/L threshold was selected as the boundary because that concentration historically has been considered the separation between the freshwater and saline-water zones of the aquifer and is a well-documented marker within the aquifer.

The potentiometric-surface map constructed from data collected in fall 2000 in a multi-agency effort organized by the Edwards Aquifer Authority indicates that water within the SCFP flows from southwest to northeast from areas of higher water-level altitude to areas of lower water-level altitude (fig. 7). Recharge to the SCFP is primarily regional, occurring in the Edwards aquifer recharge zone west of the study area. The Comal Springs fault fails to offset the entire thickness of the Edwards aquifer near the area where the boundaries of Bexar, Comal, and Guadalupe Counties come into contact. Maclay and Land (1988) referred to this area as the Bracken gap (fig. 6). Subsurface inflow to the SCFP can occur as ground water spills over from north of Comal Springs fault across the Bracken gap. The permeable sections of the aquifer in the SCFP, although faulted, are juxtaposed in a manner that promotes the flow of ground water in the study area from the center of Bexar County through southern Comal County. The ground water then encounters a transverse fault northeast of New Braunfels (fig. 6) that forms a barrier to ground-water flow and is forced up along the Comal Springs fault and through the overlying gravels to form a majority of the springs and seeps in the Comal Springs complex. The SCFP gradually narrows as it nears Comal Springs because saline water encroaches as the freshwater is discharged at the springs. The transmissivity in the freshwater zone of the Edwards aquifer southeast of the Comal Springs fault in southern Comal County is greater than in any other part of the aquifer (Maclay and Small, 1984).

Water-level hydrographs for wells Bexar, Bracken, Solms, and NBU–LCRA and the discharge hydrograph for Comal Springs (fig. 8) provided evidence of patterns of ground-water flow in the SCFP. The hydrographs for the four wells and Comal Springs fluctuated in a similar manner, possibly indicating a common flow path for ground water flowing past the wells and discharging at Comal Springs. Wells Bexar and Bracken showed larger magnitudes of fluctuation than did wells Solms and NBU–LCRA. This indicates that amplitude of the pressure wave moving through the confined section of the aquifer likely is decreased by the substantial increase in horizontal hydraulic conductivity of the aquifer in the SCFP near Comal Springs. The relatively flat potentiometric surface near Comal Springs (fig. 7) also provides evidence of higher horizontal hydraulic conductivity at the northeastern end of the SCFP.

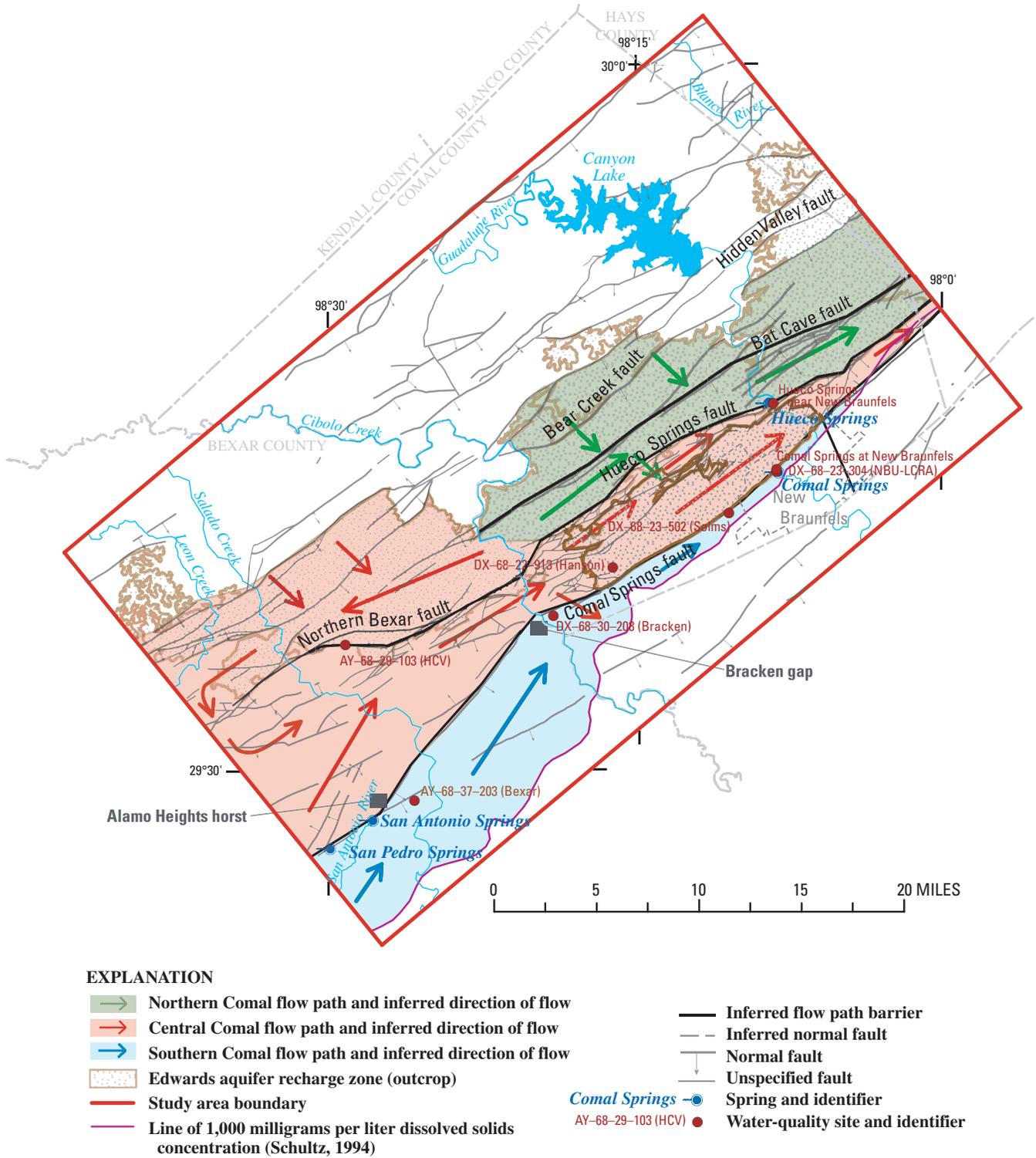


Figure 6. Locations of hypothesized ground-water flow paths, northeastern Bexar and southern Comal Counties, Texas.

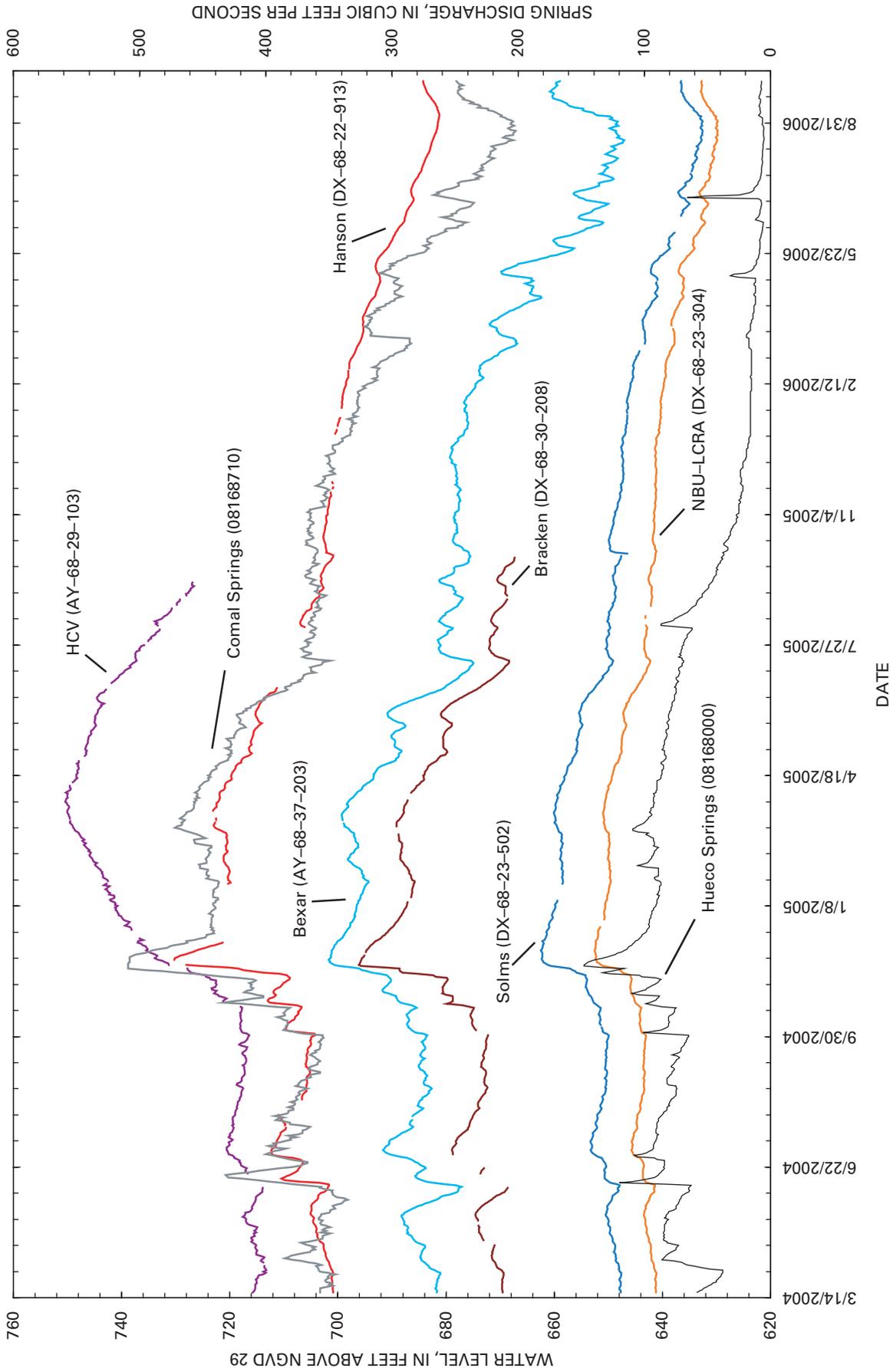


Figure 8. Hydrographs of water levels from six wells and discharge for Comal Springs and Hueco Springs, northeastern Bexar and southern Comal Counties, Texas, March 2004–September 2006.

Central Comal Flow Path

The central Comal flow path (CCFP) (fig. 6) is bounded on the northwest by the Hueco Springs fault and on the southeast by the Comal Springs fault. The potentiometric-surface map indicates that water within the CCFP flows from southwest to northeast (fig. 7). As water moves into central Bexar County from the west it is bifurcated into two flow paths by the Alamo Heights horst (Maclay and Land, 1988) (fig. 6) in the subsurface southwest of the recharge zone in central Bexar County. Water that flows north of the horst becomes constrained northeast of the Comal Springs fault, which completely offsets the Edwards aquifer east of the horst, and becomes flow within the CCFP, whereas water that flows south of the horst becomes flow within the SCFP. Local recharge to the CCFP occurs in northeastern Bexar County as water infiltrates into the Edwards aquifer outcrop northwest of the Northern Bexar fault (fig. 6). The water then moves through the subsurface to the southeast until it encounters the Northern Bexar fault. The Northern Bexar fault acts as a barrier to flow at its northeastern end (Small, 1986, fig. 8), causing water to move to the southwest along the fault trace. The displacement along the Northern Bexar fault decreases southwestward along its trace until it is no longer a barrier to flow (Small, 1986, fig. 9), which allows water to enter the confined section of the aquifer. The local recharge remains north of Comal Springs fault as it merges with the northeastward-moving water of the horst-induced CCFP. Additional inflow to the CCFP might occur in the subsurface when water from north of the Hueco Springs fault in Comal County crosses over the Hueco Springs fault in areas where the fault fails to fully offset the Edwards aquifer (Small, 1986, figs. 4–5).

As the water in the CCFP continues flowing toward the northeast, structural features in the Edwards aquifer affect the patterns of flow. In eastern Bexar County, the Bracken gap conveys water from the CCFP into the SCFP (fig. 6). Beyond the Bracken gap, the aquifer containing the CCFP is heavily faulted and partially unconfined in the area directly northwest of Comal Springs. In the confined section of the CCFP in this area, the flow is funneled into two hydraulically connected troughs that follow roughly parallel grabens formed in the fault blocks between Comal Springs fault and Hueco Springs fault (fig. 4). The southeastern trough is unconfined at its northeastern end and receives recharge from direct infiltration of rainfall and streamflow losses. The northwestern trough receives subsurface inflow across sections of the Hueco Springs fault that fail to offset the entire thickness of the Edwards aquifer. Although flow within this section of the CCFP is complicated by the complexity of the faulting, in general, much of the water in the northwestern trough likely discharges at Hueco Springs and most of the water in the southeastern trough and in the unconfined sections of the aquifer in the CCFP discharges at springs Comal 1, Comal 2, and Comal 3 at the Comal Springs complex. Water that is not discharged at either Hueco or Comal Springs continues north-eastward toward the San Marcos Springs (fig. 1).

Comparison of the water-level hydrographs for wells HCV and Hanson and discharge hydrographs for Comal Springs and Hueco Springs (fig. 8) provide insight into the movement of ground water in the CCFP. The hydrograph for well HCV shows little, if any, similarity to the hydrographs for well Hanson or either of the springs, indicating a lack of shared flow paths between well HCV and the other sites in the CCFP. Well HCV is in a heavily faulted section of the CCFP that is most likely locally isolated from the rest of the flow path.

Well Hanson is in the unconfined section of the southeastern trough of the CCFP. In general, the hydrograph for well Hanson followed the fluctuation patterns of the discharge hydrographs of Comal Springs and Hueco Springs in number and relative vertical displacement for water levels greater than about 700 feet above NGVD 29. In this area, the regional dense member of the Person Formation within the Edwards aquifer acts as a confining unit. Well Hanson is completed below the regional dense member (approximate altitude 700 feet above NGVD 29). When water levels in well Hanson rose above the regional dense member, the large magnitude of fluctuations in the hydrograph reflected confined properties similar to those at Comal Springs and Hueco Springs. When water levels in well Hanson fell to levels near or below the regional dense member, the Hanson well hydrograph tended to lose much of the vertical displacement and actually resembled the hydrographs of wells Solms and NBU–LCRA. As water levels fell in well Hanson and thus in that section of the CCFP, conditions changed from confined to unconfined. Furthermore, confined flow from the SCFP, under extreme pressure, might have pushed upward across Bracken gap into the CCFP to locally influence water levels in the CCFP.

The Hueco Springs hydrograph (fig. 8) provided evidence that the CCFP might not be the only source of water to the Hueco Springs. An increase in discharge at Hueco Springs in August 2005 had no corresponding increase at Comal Springs, indicating that the pulse of water causing the increased discharge at Hueco Springs either was not sourced in the CCFP or did not reach Comal Springs.

Northern Comal Flow Path

The northern Comal flow path (NCFP) (fig. 6) is in the Edwards aquifer recharge zone north of Hueco Springs fault and south of Bat Cave fault in Comal County. The potentiometric-surface map indicates that water in the NCFP flows from southwest to northeast (fig. 7). Recharge to the NCFP occurs from direct infiltration of rainfall and streamflow losses to the Edwards aquifer exposed at the surface in the NCFP. Additional inflow likely comes from the Trinity aquifer in the subsurface where the Bat Cave fault juxtaposes the Trinity aquifer against the Edwards aquifer. An undetermined amount of water from the NCFP might flow into the confined sections of the CCFP across Hueco Springs fault in areas where the fault does not completely offset the Edwards aquifer. The

water remaining north of the Hueco Springs fault flows toward San Marcos Springs.

Correlation of Water Levels

Statistical correlation between water levels for the six observation wells and spring discharges for Comal Springs and Hueco Springs provides additional evidence for the hypothesized flow paths. Figure 9 graphically demonstrates the relations involving each of the 28 combinations of paired datasets. The Pearson's *r* correlation coefficient indicates the strength of the linear correlation between dataset pairs and, thus, between water levels or spring discharges, or both. The larger the absolute value of *r* between -1 and 1, the stronger the linear correlation. The strongest correlations (*r*-values ranging from .991 to .999) were observed between the datasets for wells NBU-LCRA and Solms (SCFP) and well Hanson (CCFP), providing evidence for the hydraulic connection between the SCFP and the CCFP across the Bracken gap. The dataset for well Bracken, the well closest to the Bracken gap, is not as strongly correlated with datasets for wells NBU-LCRA, Solms, and Hanson. The well Bracken might be in a less transmissive part of the aquifer than the other three wells, allowing for greater fluctuations in water levels in the well Bracken dataset, which could account for weaker correlations with datasets from those wells. The dataset for well HCV displayed the weakest correlations with the other datasets, indicating a lack of shared flow paths with water at the other wells, Comal Springs, or Hueco Springs. The dataset for Comal Springs was strongly correlated with those for wells Bexar, NBU-LCRA, Hanson, and Solms (*r*-values ranging from .961 to .983) and fairly strongly correlated with the dataset from well Bracken (*r*-value of .954). The dataset for Hueco Springs did not correlate linearly with any of the datasets from the other wells or Comal Springs but did have monotonic relations (an increase in one variable corresponds to an increase in the other) with the datasets from wells Bexar, NBU-LCRA, Hanson, and Solms and Comal Springs. One reason for the lack of linear correlation between Hueco Springs and the other Edwards aquifer sites could be the influence of inflow from the Trinity aquifer. In addition, the northwestern trough of the CCFP, which provides most of the flow from the Edwards aquifer to Hueco Springs, could be less hydraulically connected to the rest of the CCFP during periods of low spring discharge than during periods of high spring discharge.

Analysis of Hydrograph Recession Curves

Hydrographs of discharge from Comal Springs and Hueco Springs were analyzed. Recession curves representing a range of flows were selected from each hydrograph for use in the analyses of hydrograph recession curves. The recession curves were examined for breaks (inflection points) in the recession slope, which are indicative of a change from one flow regime to another within the karst continuum. A recession

coefficient (α) was calculated for each section of hydrograph between inflection points using Milanovich's (1981) equation,

$$Q_t = Q_0 e^{\alpha(t-t_0)}$$

where Q_0 is the initial discharge at time t_0 , the beginning of each recession slope, and Q_t is the discharge at time t . Solving for α yields the equation

$$\alpha = \frac{\ln(Q_t/Q_0)}{(t-t_0)}$$

For the hydrographs of this study (fig. 9), an α of about 0.18, corresponding to a steep (nearly vertical) slope, indicates conduit-driven drainage of a karst aquifer. An α of about 0.008, corresponding to a more horizontal slope, indicates the diffuse drainage of the primary porosity in the matrix of the aquifer. An α of about 0.02–0.09, corresponding to an intermediate slope, can indicate either drainage of the fractures in the aquifer or a mixture of conduit and diffuse drainage.

Figure 10 illustrates discharge magnitude and variability for Comal Springs and Hueco Springs from mid-March 2004 through August 2006. Comal Springs discharge was consistently greater and more variable than discharge at Hueco Springs. In general, increases in discharge at Hueco Springs were smaller and shorter in duration than increases at Comal Springs.

Discharge-recession curves for Comal Springs were used to identify the flow regimes in the aquifer contributing to the springs. Flow varied from 202 to 509 cubic feet per second (ft³/s) during 2004–06 with the highest discharges occurring during November 21–25, 2004. The recession curve for the highest-discharge period indicates that fracture flow, or a mixture of conduit flow and diffuse flow, first dominated the ground-water system supplying water to the springs, followed by conduit flow and then fracture flow, or a mixture of conduit and diffuse flow, again. This pattern indicates that, during periods of high-discharge recession at the Comal Springs complex, the ground-water system supplying the springs might include upper and lower layers of the aquifer dominated by fractures or a mixture of conduits and aquifer matrix, with a middle section of the aquifer dominated by conduits. The lowest discharges occurred near the end of the study period (August 2006). During the lowest-discharge period, July 2005–August 2006, the recession curves for Comal Springs discharge reflected mostly an α of about 0.09, indicating fracture flow or a mixture of conduit and diffuse flow. The majority of the recession slopes for the Comal Springs hydrograph for the 2-year period had an α of about 0.09, indicating that the overall flow through the aquifer to Comal Springs is fracture flow or a mixture of conduit and diffuse flow. During the 2-year period, α values less than 0.02 were not observed for Comal Springs, indicating that diffuse flow never dominated the flow system supplying Comal Springs.

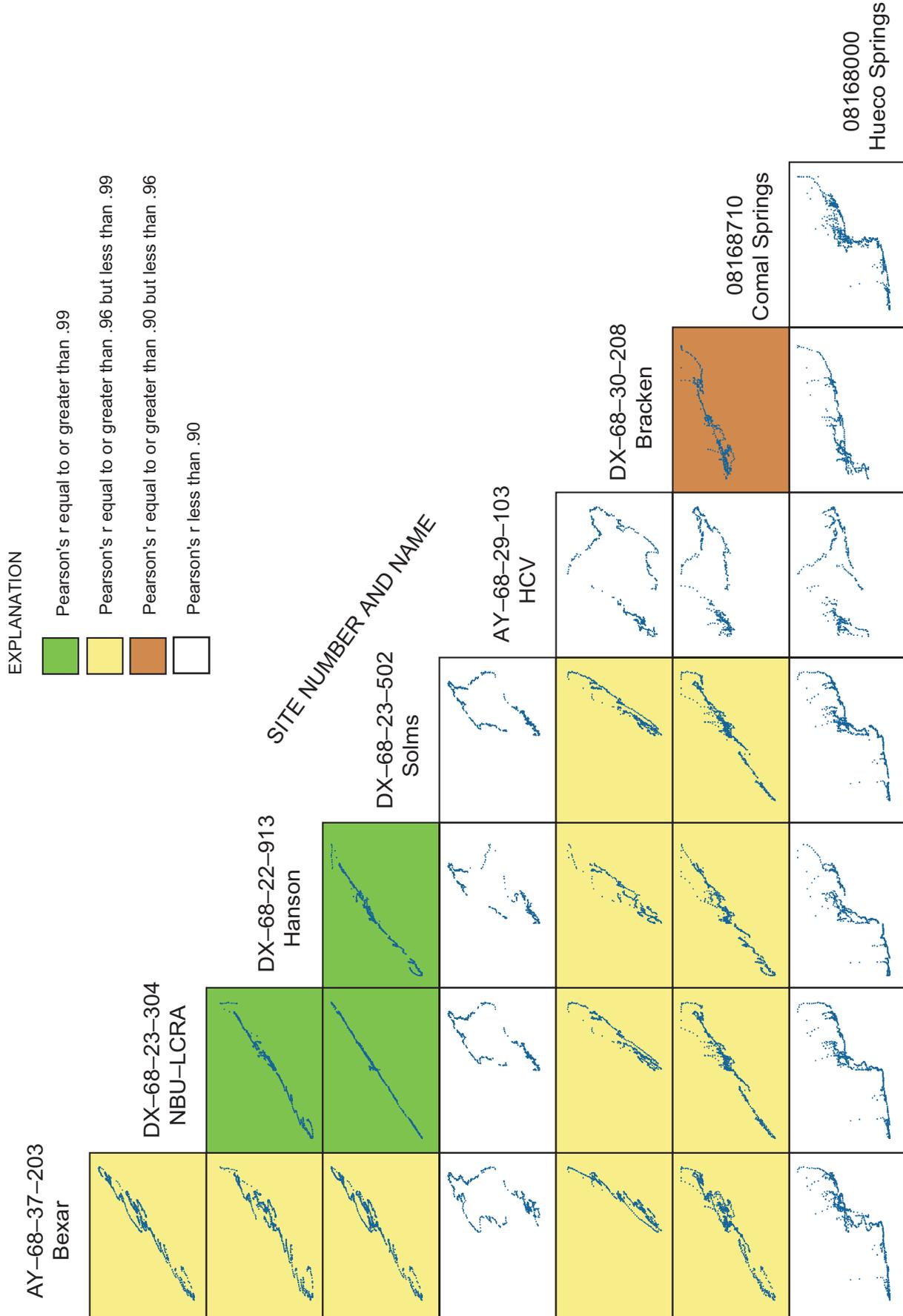


Figure 9. Correlations between water levels for six wells and discharge for Comal Springs and Hueco Springs, northeastern Bexar and southern Comal Counties, Texas, March 2004–September 2006.

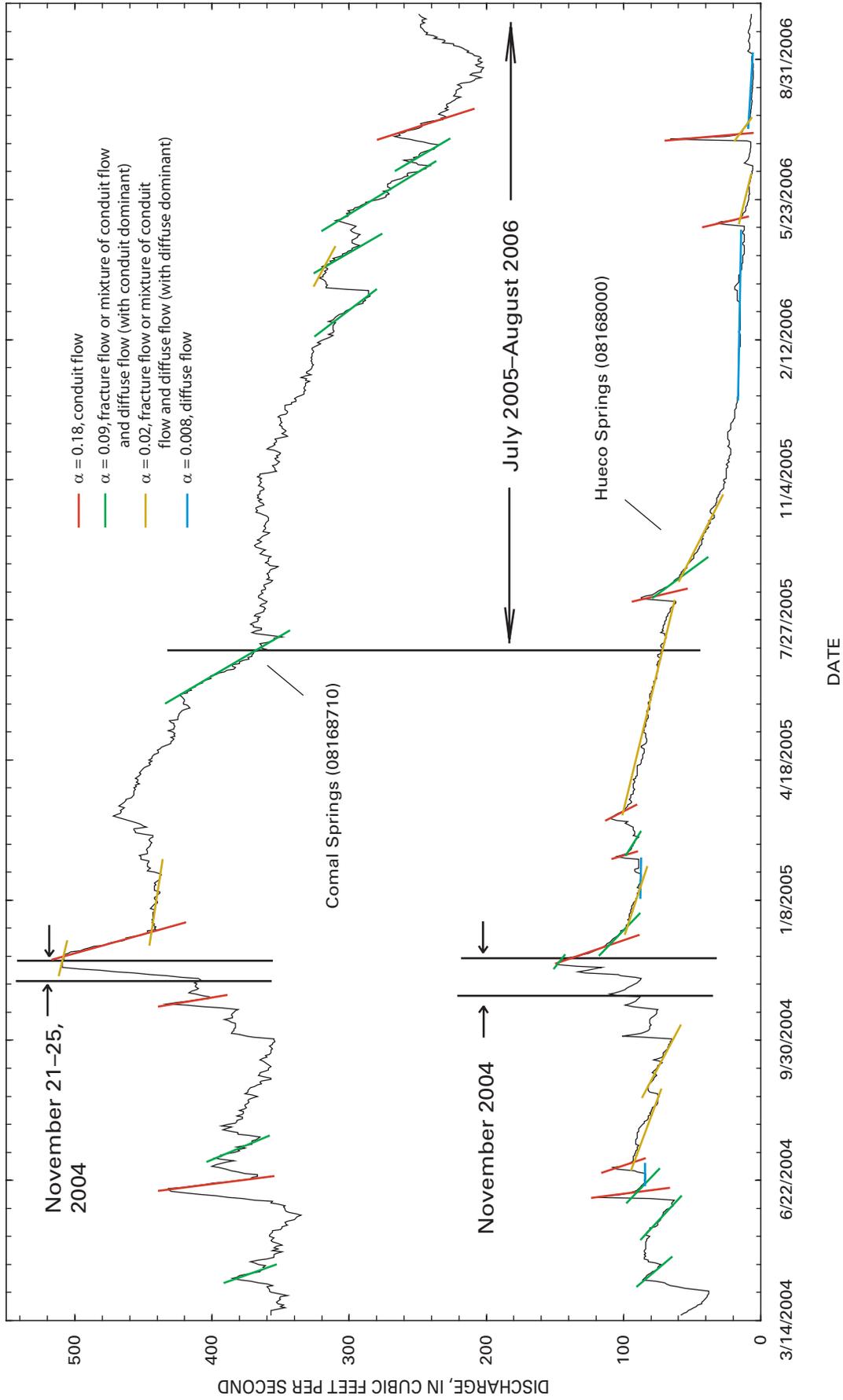


Figure 10. Hydrographs of discharge at Comal Springs and Hueco Springs, southern Comal County, Texas, March 2004–September 2006, showing recession coefficients for recession curves.

Discharge-recession curves for Hueco Springs were analyzed to identify the flow regimes in the aquifer contributing to the springs and to attempt to obtain additional evidence of contribution from the Trinity aquifer. Flow at Hueco Springs varied from 5.3 to 148 ft³/s during 2004–06 with the highest discharges occurring during November 2004. The recession curve for the highest-discharge period indicates that fracture flow or a mixture of conduit and diffuse flow first dominated the ground-water system supplying water to the springs. Following this period of fracture or mixed drainage, conduit flow controlled the release of water to the springs, followed by another period of fracture or mixed flow. The second period of fracture or mixed drainage was followed by periods in which diffuse flow became increasingly dominant and eventually prevailed. The lowest discharges occurred near the end of the study period (August 2006). During the lowest-discharge period, July 2005–August 2006, the increases in discharge at Hueco Springs were characterized by sudden high increases with recession curves that were dominated by conduit flow followed by fracture or conduit/diffuse mixed flow and ending with diffuse flow. During 2004–06, the recession slopes with an α of about 0.18 represented the greatest number of slopes for the Hueco Springs hydrograph, indicating that initial drainage of the aquifer to Hueco Springs following a recharge event is dominated by a fast-draining conduit system. Following the drainage of the conduits, the system typically is characterized by either fracture flow or a mixture of conduit and diffuse flow, or diffuse flow.

Although the Comal Springs complex and Hueco Springs are only about 3 miles apart, the hydrograph recession slope analysis for the two springs reinforces the fact that the springs are sourced from a complex karst system. The initial recession slopes for Comal Springs predominately represent a mixture of conduit and fracture flow, or a mixture of conduit and diffuse flow, whereas the initial recession slopes for Hueco Springs predominately represent conduit flow. The initial mixed flow from Comal Springs likely represents the conduit flow contribution from the SFCF, which provides most of the flow from the springs, and the fracture flow contribution from the CCFP. The initial conduit flow from Hueco Springs provides evidence that the northwestern trough within the CCFP in Comal County might be highly transmissive, allowing for rapid movement of water through the aquifer to Hueco Springs.

Geochemical Identification of Flow Paths

Major-Ion Chemistry

Trilinear diagrams constructed using major-ion chemistry from samples collected at wells in and springs emerging from the SFCF, CCFP, and NCFP (fig. 11) indicate that all

three flow paths are dominated by calcium-bicarbonate type water. Relatively little variation in major-ion chemistry was observed. The water samples from wells and springs in the CCFP showed more variability in major-ion chemistry than those from the two other flow paths. The samples from wells in the NCFP consistently had higher percentages of bicarbonate and lower percentages of sulfate and chloride than the samples collected from sites in the SFCF. The comparatively high sulfate and chloride concentrations in the SFCF samples might indicate that the water in the flow path was mixing with small amounts of saline water from the freshwater/saline-water transition zone.

A graph of the relation between calcium concentration and the ratio of magnesium to calcium concentrations provided further evidence for the hypothesized flow paths (fig. 12). Samples from wells in the SFCF generally yielded low calcium concentrations coupled with moderate Mg/Ca ratios (thus moderate magnesium concentrations). The highest calcium concentrations and lowest Mg/Ca ratios (contributed to by low magnesium concentrations) were in samples collected from the CCFP. The samples from the CCFP yielded data points that plotted in a grouping overlapping the data points for the SFCF. The high concentrations of magnesium in the samples collected from the recharge zone in Comal County (NCFP) might be evidence that water from the Trinity aquifer is entering the Edwards aquifer in the subsurface. The calcium concentration of water from Hueco Springs was lower when the spring discharge was low than when the spring discharge was high, indicating that during periods of low discharge Hueco Springs might receive a larger contribution from the Trinity aquifer than during periods of high discharge.

Stable Isotopes

The relation between δD and $\delta^{18}O$ in water samples from the Edwards aquifer, by flow path, is shown in figure 13. The graph shows that, except for samples collected from wells in the Edwards aquifer recharge zone in northeastern Bexar County (CCFP) following an unusually intense rain storm produced by a tropical system in October 1998, there was not much variation in stable isotope values among the flow paths. In the study area δD in ground water ranged from -36.00 per mil (well AY-68-28-313 in the CCFP) to -20.89 per mil (well DX-68-23-620 in the SFCF) and $\delta^{18}O$ ranged from -6.03 per mil (well AY-68-28-313 in the CCFP) to -3.70 per mil (spring Hueco A in the CCFP). Excluding ground-water samples collected from wells in the recharge zone of northeastern Bexar County following the October 1998 storm, the δD range in the study area was -33.00 to -20.89 per mil and the $\delta^{18}O$ range was -4.60 to -3.70 per mil.

The local meteoric water line (LMWL; $\delta D=8.8032 \delta^{18}O+17.825$) was calculated using rainfall isotope data collected in the study area. All ground-water samples, except one, plotted below the LMWL. The plotting positions indicate that the water evaporated, to varying degrees, before entering the aquifer as recharge.

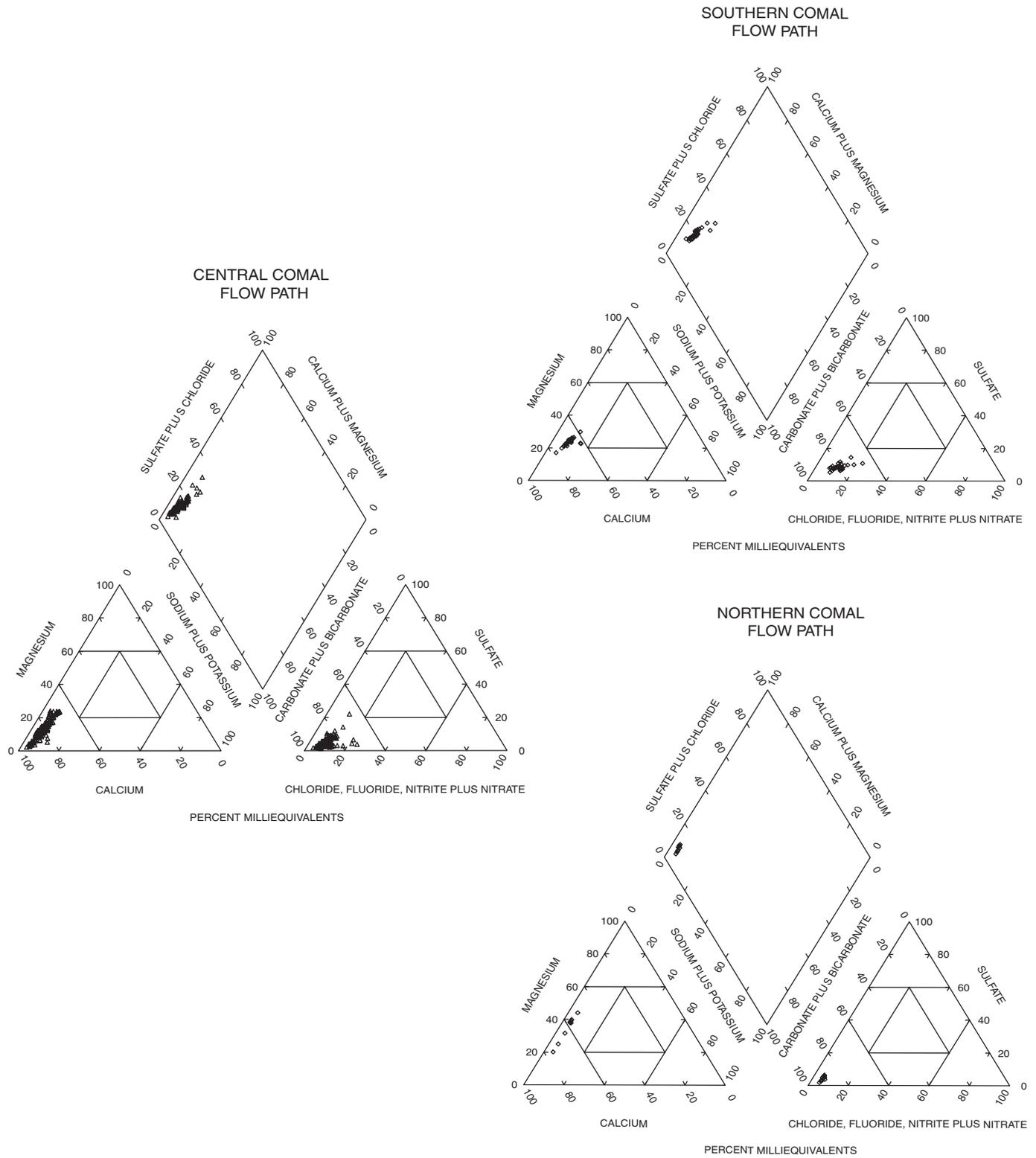


Figure 11. Ternary diagrams showing composition of ground water in the Edwards aquifer in the southern, central, and northern Comal flow paths, northeastern Bexar and southern Comal Counties, Texas.

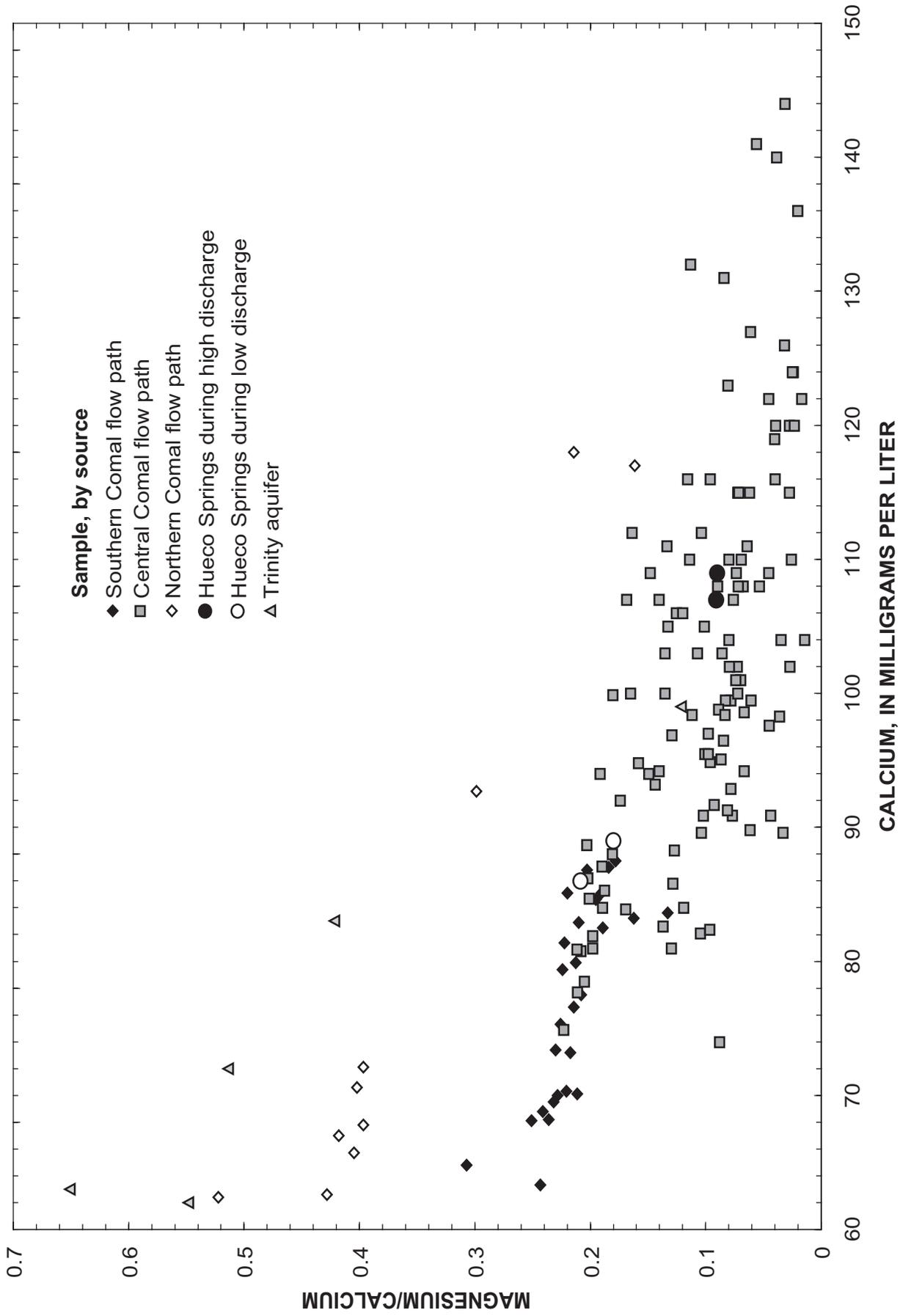


Figure 12. Relation of calcium concentration to magnesium/calcium ratio in samples collected from wells and springs in the Edwards aquifer and wells in the Trinity aquifer, northeastern Bexar and southern Comal Counties, Texas.

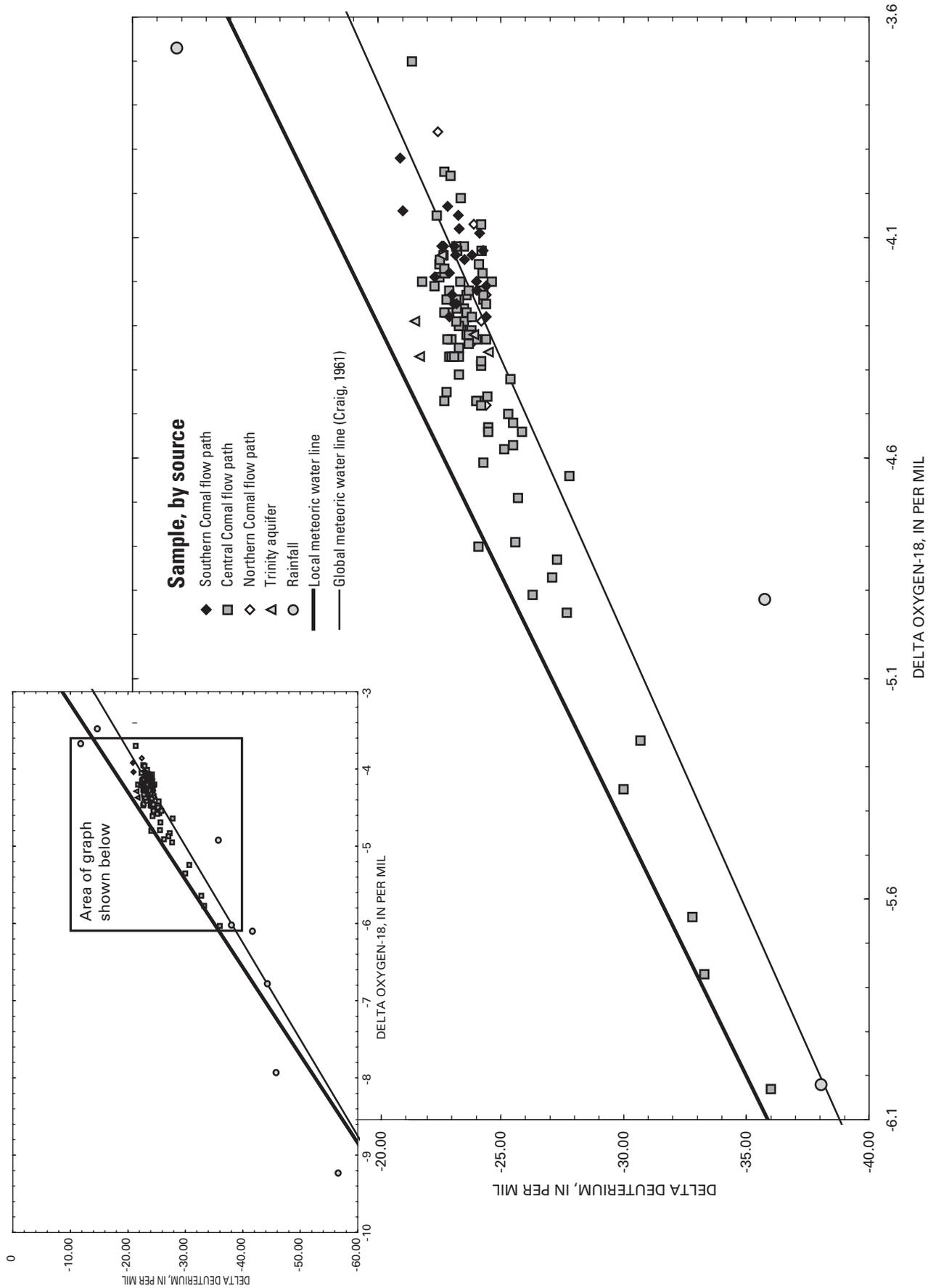


Figure 13. Relation between delta oxygen-18 and delta deuterium for selected wells and springs, northeastern Bexar and southern Comal Counties, Texas, 1996–2006, with insert showing same plot with axes extended to range of rainfall samples.

Samples collected from wells and springs in the SCFP plotted in a relatively small area on the graph (fig. 13) with δD values ranging from -24.40 to -20.89 per mil and $\delta^{18}O$ values ranging from -4.28 to -3.92 per mil. The largest range in isotopic composition was observed in samples collected from the CCFP, with δD values ranging from -36.00 to -21.39 per mil and $\delta^{18}O$ values ranging from -6.03 to -3.70 per mil. CCFP samples with δD values greater than about -4.3 per mil overlap the grouping of SCFP samples. The samples from wells in the NCFP plot in a narrow area that spans the area occupied by the SCFP samples with δD values ranging from -24.40 to -22.44 per mil and $\delta^{18}O$ values ranging from -4.48 to -3.86 per mil.

Ground-Water Ages

Apparent ages of ground water in samples collected in 2003 derived from SF_6 analysis (fig. 14) ranged from about 21 years in well DX-68-23-104 in the NCFP to present day (less than 0.5 year or recharged in 2003) at spring Comal 7 in the SCFP. Ages ranged from about 17 years (well AY-68-29-914) to present day (spring Comal 7) in the SCFP, with the oldest water in the upgradient section of the SCFP and the youngest water from a spring at the downgradient end of the SCFP. The young apparent age of the Comal 7 spring might be evidence that young water from the upthrown block of Comal Springs fault is supplying additional recharge to the spring. The SF_6 method of age dating water does not take into account the mixing of young and old waters and, therefore, reflects the age of the young water. In the CCFP, ages ranged from about 16 years (well DX-68-22-810) to about 1 year (well DX-68-30-111), with both the oldest and youngest water in the upgradient section of the flow path. Water in the NCFP likely receives inflow from the subsurface as a result of inter-aquifer hydraulic connection with the Trinity aquifer, which has substantially lower transmissivity (thus, possibly water of older apparent age) than does the Edwards aquifer. The apparent age of the water in well DX-68-23-104 might reflect the influence of the Trinity aquifer on the NCFP.

The apparent ages of the samples collected in 2003 from the various orifices of the Comal Springs complex and Hueco Springs, computed using SF_6 analysis, provide evidence of the complex nature of the karst aquifer in which the springs are sourced. The samples from the orifices at the Comal Springs complex range in age from about 6 years to present day (less than 0.5 year or recharged in 2003). The samples from Hueco Springs provided ages that range from about 7 to 8 years. The recent ages observed from the orifices at Comal Springs indicate an inflow of water near the springs complex, possibly from the unconfined section of the Edwards aquifer at the northeastern end of the southeastern trough of the CCFP in Comaal County. The relatively old water at Hueco Springs indicates a possible inflow of water into the ground-water

system supplying the springs from a source of water yielding relatively older water than the Edwards aquifer, possibly the Trinity aquifer.

Apparent ages of ground-water samples collected in 2003 derived from $^3H/^3He$ analysis (fig. 15) ranged from greater than 45 years in well DX-68-23-104 in the NCFP to present day (less than 0.5 year or recharged in 2003) in well DX-68-22-810 in the CCFP. Ages ranged from 21.3 to 8.8 years in the SCFP. In general, the samples from wells in Bexar County (upgradient section of the SCFP) showed younger ages than wells and springs in the downgradient section of the SCFP. The samples from a spring in the Comal Spring complex, Spring Island, indicated ages similar to those for samples from the wells in Bexar County. A possible explanation for the anomalously young age of water in the Comal Springs complex is that there is a local source for the spring. Well DX-68-23-602 in the SCFP had a younger age than the wells in Bexar County. This well is on the escarpment formed along the Comal Springs fault and might be receiving inflow from the unconfined CCFP on the upthrown side of the fault. In the CCFP, ages ranged from 17.3 years to present day (less than 0.5 year or recharged in 2003) with the oldest water in well AY-68-30-520 near the Bexar-Comal County line in the central section of the flow path and the youngest water in the central section of the flow path where the Edwards aquifer is exposed, allowing recharge from infiltration of rainfall and streamflow losses to occur. The data from samples collected at Hueco Springs indicate very young water (apparent age less than 0.5 year) mixed with very old water. The oldest water was in well DX-68-23-104 in the NCFP. Water in this well is tritium depleted, indicating that the water has undergone complete tritium decay. The apparent age of the water is not attainable by the $^3H/^3He$ method, but based on the tritium decay rate the water is estimated to have been recharged more than 45 years ago. A possible source for this old water could be inflow from the underlying Trinity aquifer.

Apparent ages for ground-water samples collected in 2006 ranged from 41.1 years at spring Comal 7 in the SCFP to present day (less than 0.5 year or recharged in 2006) in wells DX-68-23-710 and DX-68-22-905 in the CCFP (fig. 15). In general, samples collected in 2003 yielded younger apparent ages than samples collected during 2006. Water levels in the Edwards aquifer were higher during the sampling in 2003 than they were during the sampling period in 2006, likely resulting in faster flow rates in the aquifer because of increased hydraulic gradients. Many of the samples contained a mixture of young and old water, especially those collected from areas near the downgradient ends of the flow paths. The samples collected from Hueco Springs in 2006, like those collected in 2003, showed a mixture of very young water (apparent age less than 0.5 year) with very old water (possibly sourced from the Trinity aquifer).

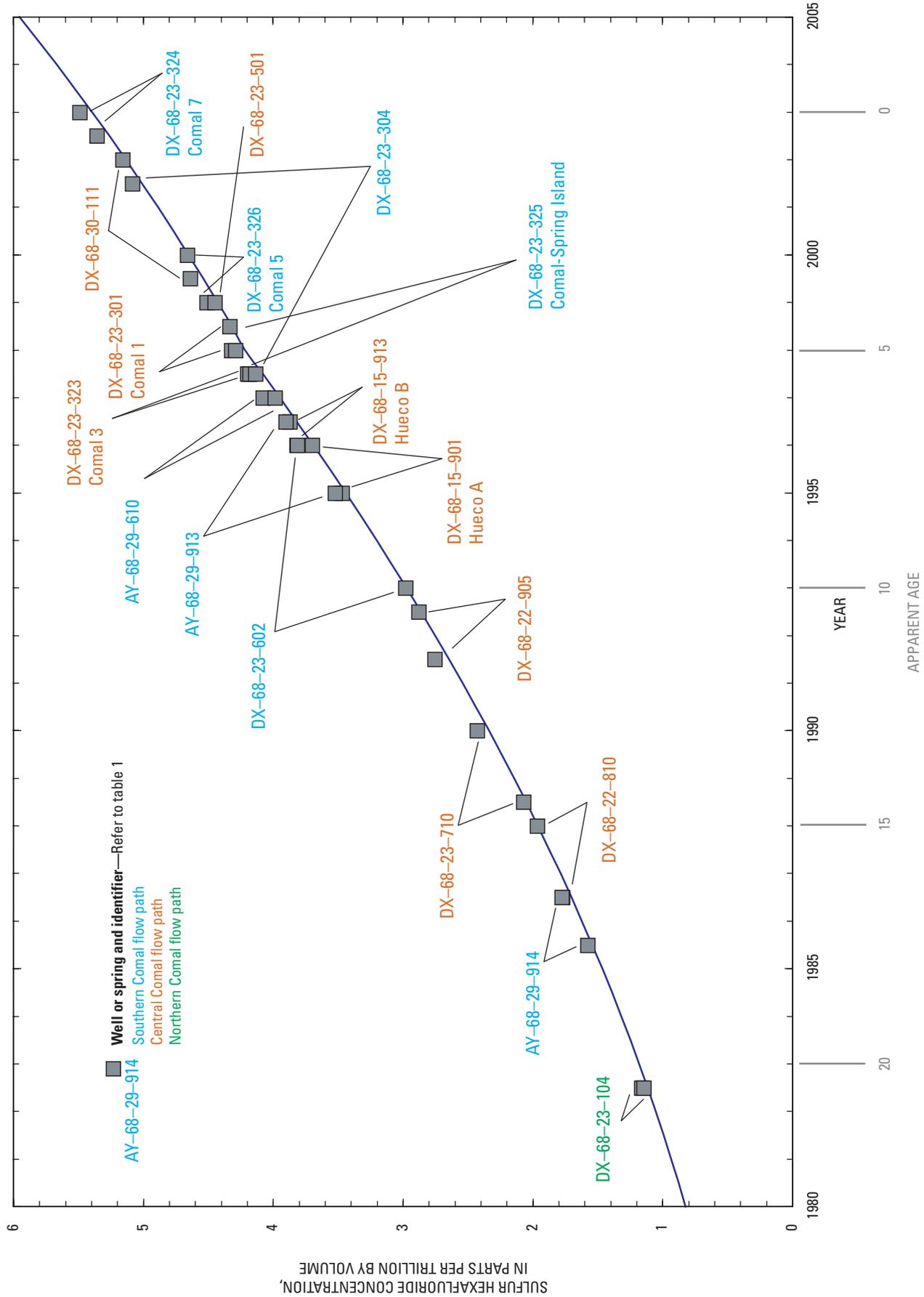
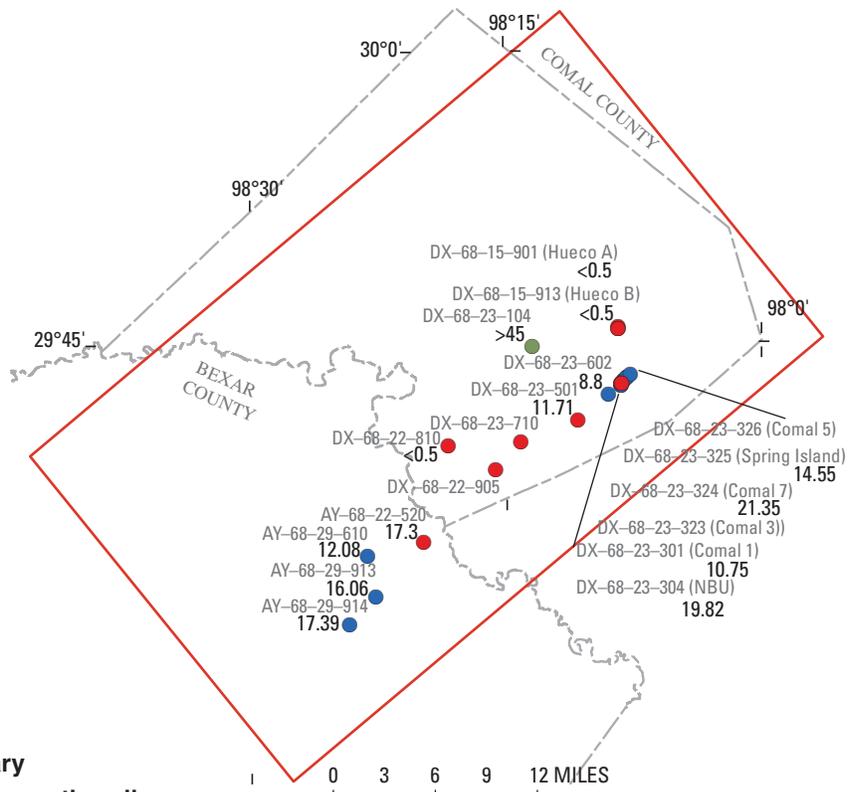


Figure 14. Apparent ages of ground water in 2003 samples from selected wells and springs in the Edwards aquifer, northeastern Bexar and southern Comal Counties, Texas, based on sulfur hexafluoride concentration (modified from Busenberg and Plummer, 2000, fig. 6A).

Sample year 2003

EXPLANATION

- Study area boundary
- AY-68-29-913 16.06 Southern Comal flow path well, identifier, and age of water in years
- DX-68-23-501 11.71 Central Comal flow path well, identifier, and age of water in years
- DX-68-23-104 >45 Northern Comal flow path well, identifier, and age of water in years
- > = greater than
- < = less than



Sample year 2006

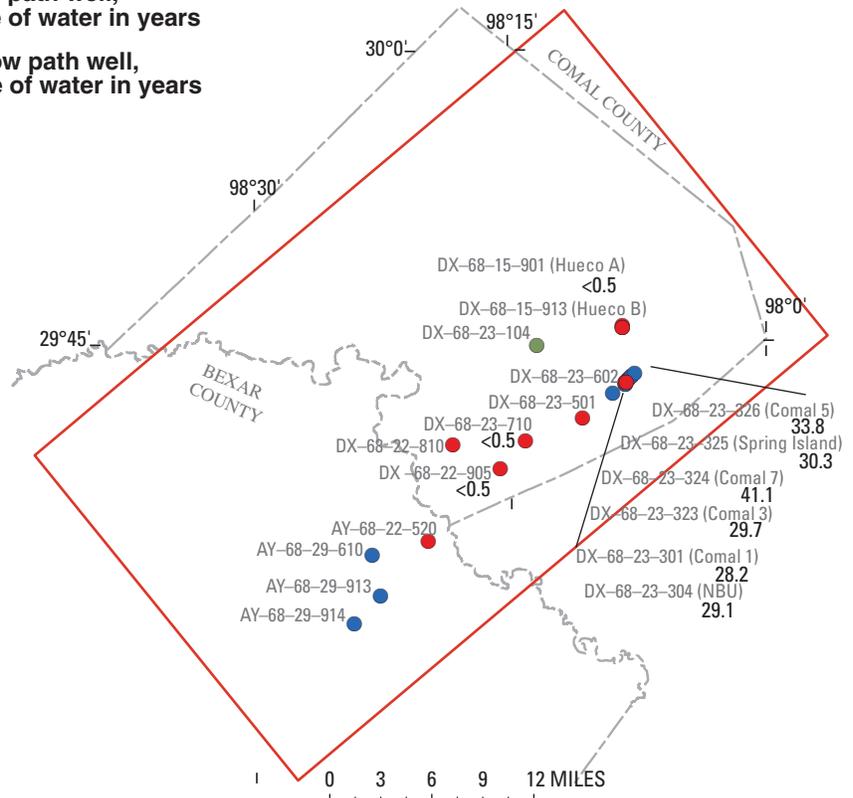


Figure 15. Apparent age of water in selected wells in the Edwards aquifer, northeastern Bexar and southern Comal Counties, Texas, based on tritium/helium-3 ratio.

Summary

The U.S. Geological Survey (USGS), in cooperation with the San Antonio Water System, conducted a 4-year study during 2002–06 to identify major flow paths in the Edwards aquifer in northeastern Bexar and southern Comal Counties (study area). Geologic structure, surface- and ground-water data, and geochemistry were used to identify flow paths. Knowledge of geologic structure and previous flow path analyses conducted by the USGS provided a basis for the initial selection of hypothesized flow paths in the study area.

Historical and new data were used in analyses. Continuous water-level data were collected at six observation wells from mid-March 2004 through September 2006. Discharge data were obtained from two springflow sites for the same time period. Ground-water chemistry and isotope data were compiled from samples collected from 76 wells and nine springs (and spring orifices of major springs) during 1996–2006. Isotope data also were obtained from samples collected at four rainfall-sampling sites during 1998–2003.

The southern Comal flow path (SCFP) is bounded on the northwest by the Comal Springs fault and on the southeast by the freshwater/saline-water transition zone in the Edwards aquifer. Most of the water in this flow path enters the aquifer west of the study area. Additional inflows occur as spillover from north of the Comal Springs fault at the Bracken gap, an area near the Bexar-Comal-Guadalupe County line intersection where the fault has not completely offset the entire aquifer thickness. The SCFP gradually narrows as it nears Comal Springs because saline water encroaches as the freshwater is discharged at the springs.

The central Comal flow path (CCFP) is bounded on the northwest by the Hueco Springs fault and on the southeast by the Comal Springs fault. Water that is diverted north by the Alamo Heights horst in central Bexar County flows north of Comal Springs fault, accounting for the regional source to this flow path. Local recharge from infiltration of rainfall and streamflow leakage is supplied to the flow path from the Edwards aquifer recharge zone in northeastern Bexar County. These mixed waters then flow northeast through a highly fractured and faulted section of the aquifer. The complex faulting produces a hydraulically connected, two-trough system in the aquifer that is confined in some areas and unconfined in others. The troughs channel flow to discharge points at both Comal Springs and Hueco Springs. Water that does not discharge at either of the springs continues flowing northeastward toward San Marcos Springs.

The northern Comal flow path (NCFP) is in the Edwards aquifer recharge zone in Comal County and is bounded on the north by the Bat Cave fault and on the south by the Hueco Springs fault. Recharge to this flow path occurs from the Edwards aquifer recharge zone northwest of Bat Cave fault in Comal County from direct infiltration of rainfall and streamflow losses to the Edwards aquifer exposed at the surface. Additional inflow likely comes from the Trinity aquifer in the

subsurface where the Bat Cave fault juxtaposes the Trinity aquifer against the Edwards aquifer. Flow also might occur between the CCFP and the NCFP in central Comal County through a section of the Hueco Springs fault that does not completely offset the entire thickness of the Edwards aquifer.

A potentiometric-surface map derived from synoptic water-level measurements made in fall 2000 was used to identify the generally southwest to northeast flow directions within the flow paths. Statistical correlations between water levels for six observation wells and discharges from Comal Springs and Hueco Springs (28 combinations of paired datasets) yielded additional evidence for the hypothesized flow paths. Strong linear correlations were evident between the datasets from wells and springs within the same hypothesized flow path and the datasets from wells in areas where flow between flow paths was suspected.

Hydrograph recession slope analysis for Comal and Hueco Springs reinforces the fact that the springs are sourced from a complex karst system. The initial recession slopes for Comal Springs predominately represent a mixture of conduit and fracture flow, or a mixture of conduit and diffuse flow, whereas the initial recession slopes for Hueco Springs predominately represent conduit flow. The initial mixed flow from Comal Springs likely represents the conduit flow contribution from the SCFP, which provides most of the flow from the springs, and the fracture flow contribution from the CCFP. The initial conduit flow from Hueco Springs provides evidence that the northwestern trough within the CCFP in Comal County might be highly transmissive, allowing for rapid movement of water through the aquifer to Hueco Springs.

Geochemical data (major ions, stable isotopes, sulfur hexafluoride, and tritium and helium) were used in graphical analyses to obtain evidence of the flow path from which wells or springs derive water. Major-ion geochemistry in samples from selected wells and springs showed relatively little variation. Samples from the SCFP were characterized by relatively high sulfate and chloride concentrations possibly indicating that the water in the flow path was mixing with small amounts of saline water from the freshwater/saline-water transition zone. Samples from the CCFP yielded the most varied major-ion geochemistry of the three hypothesized flow paths. CCFP samples were characterized, in general, by high calcium concentrations and low magnesium concentrations. Samples from the NCFP were characterized by relatively low sulfate and chloride concentrations and high magnesium concentrations. The high magnesium concentrations characteristic of NCFP samples from the recharge zone in Comal County might indicate water from the Trinity aquifer is entering the Edwards aquifer in the subsurface.

A graph of the relation between the stable isotopes δD and $\delta^{18}O$ showed that, except for samples collected from wells in the Edwards aquifer recharge zone in northeastern Bexar County following an unusually intense rain storm produced by a tropical system, there was not much variation in stable isotope values among the flow paths. In the study area δD in ground water ranged from -36.00 to -20.89 per mil and $\delta^{18}O$

ranged from -6.03 to -3.70 per mil. Excluding samples collected from wells in the recharge zone of northeastern Bexar County following the intense rain storm, the δD range in the study area was -33.00 to -20.89 per mil and the $\delta^{18}O$ range was -4.60 to -3.70 per mil.

Two ground-water age-dating techniques, sulfur hexafluoride concentrations and tritium-helium-3 isotope ratios, were used to compute apparent ages (time since recharge occurred) of water samples collected from wells and springs in the study area. In general, the apparent ages computed by the two methods do not seem to indicate direction of flow. Apparent ages computed for water samples in northeastern Bexar and southern Comal Counties do not vary greatly except for some very young water in central Comal County. Additional recharge from direct infiltration of precipitation and stream-flow losses into the exposed Edwards aquifer in central Comal County might account for younger age dates in the middle section of the CCFP.

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**Appendix 1—Water-Level (2004–06) and Chemical Data
(2003–06)**

Appendix 1.1. Daily mean depth to water at well DX-68-22-913, Comal County, Texas, 2004-06.

[In feet below land surface; ---, not collected or computed]

Water year 2004 (Oct. 2003–Sept. 2004) daily mean values												
Day	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	---	---	---	---	---	---	109.05	106.24	106.52	98.11	99.03	104.68
2	---	---	---	---	---	---	109.10	106.11	106.84	97.78	---	104.87
3	---	---	---	---	---	---	109.06	105.97	107.44	97.75	---	104.86
4	---	---	---	---	---	109.37	108.94	105.95	107.85	97.66	---	104.77
5	---	---	---	---	---	109.32	108.75	106.01	107.97	97.68	---	104.63
6	---	---	---	---	---	109.30	108.64	106.05	108.07	97.85	---	104.48
7	---	---	---	---	---	109.28	108.58	105.96	108.20	98.24	---	104.15
8	---	---	---	---	---	109.23	108.51	105.86	108.34	98.50	---	104.08
9	---	---	---	---	---	109.30	108.42	105.70	108.32	98.53	---	104.05
10	---	---	---	---	---	109.36	108.38	105.53	106.48	98.58	---	104.08
11	---	---	---	---	---	109.40	108.24	105.44	102.91	98.63	---	104.23
12	---	---	---	---	---	109.42	107.99	105.42	100.11	98.73	---	104.24
13	---	---	---	---	---	109.35	107.93	105.52	99.56	98.85	103.36	104.27
14	---	---	---	---	---	109.27	107.83	105.57	99.98	98.97	103.34	104.31
15	---	---	---	---	---	109.11	107.75	105.40	100.60	99.13	103.33	104.23
16	---	---	---	---	---	109.11	107.77	105.14	101.16	99.28	103.41	104.19
17	---	---	---	---	---	109.11	107.74	104.97	101.65	99.33	103.57	104.28
18	---	---	---	---	---	109.11	107.60	104.91	102.05	99.40	103.87	104.43
19	---	---	---	---	---	109.10	107.49	104.99	102.48	99.58	104.18	104.41
20	---	---	---	---	---	109.05	107.43	105.15	102.66	99.99	104.14	104.40
21	---	---	---	---	---	109.05	107.38	105.15	102.81	100.31	104.23	104.50
22	---	---	---	---	---	109.03	107.35	105.11	103.02	100.40	104.14	104.70
23	---	---	---	---	---	108.99	107.38	105.06	103.22	---	104.16	104.90
24	---	---	---	---	---	108.94	107.35	105.12	103.39	100.20	104.22	105.06
25	---	---	---	---	---	108.94	107.06	105.28	103.56	100.26	104.29	105.04
26	---	---	---	---	---	108.96	106.73	105.55	103.41	99.98	104.33	104.92
27	---	---	---	---	---	108.90	106.53	105.83	103.34	99.81	104.30	104.88
28	---	---	---	---	---	108.83	106.38	106.12	102.89	99.66	104.24	105.01
29	---	---	---	---	---	108.82	106.25	106.31	101.68	99.41	104.18	105.07
30	---	---	---	---	---	108.88	106.19	106.32	99.86	99.21	104.22	105.10
31	---	---	---	---	---	108.99	---	106.37	---	98.99	104.46	---
Mean	---	---	---	---	---	---	107.79	105.62	103.88	---	---	104.56
Maximum	---	---	---	---	---	---	109.10	106.37	108.34	---	---	105.10
Minimum	---	---	---	---	---	---	106.19	104.91	99.56	---	---	104.05

Appendix 1.1. Daily mean depth to water at well DX-68-22-913, Comal County, Texas, 2004-06—Continued.

Water year 2005 (Oct. 2004–Sept. 2005) daily mean values												
Day	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	105.75	98.49	80.83	---	89.78	89.41	88.23	92.37	95.10	---	---	106.10
2	105.67	98.18	81.73	---	89.85	89.34	88.13	92.71	94.87	---	---	106.51
3	103.63	98.07	82.56	---	89.93	89.47	88.10	93.38	94.77	---	---	106.60
4	101.08	98.13	83.22	---	89.96	89.41	88.10	93.80	94.75	---	---	106.57
5	100.38	98.37	83.84	---	89.86	89.31	88.14	94.04	94.78	---	---	106.55
6	100.37	98.85	84.39	---	89.75	88.86	88.35	93.76	94.83	---	---	106.63
7	100.47	99.26	85.50	---	89.67	88.32	88.82	93.63	94.95	---	---	106.99
8	100.61	99.55	86.30	---	89.42	87.65	89.05	93.63	95.12	---	---	107.24
9	100.71	99.76	87.16	---	89.27	87.28	89.18	93.55	95.27	---	---	107.34
10	100.79	99.94	88.13	---	89.14	87.05	89.06	93.60	95.40	---	103.79	107.18
11	101.08	100.31	88.66	---	88.99	87.02	89.13	93.91	95.46	---	103.19	107.11
12	101.43	100.66	---	---	88.86	87.15	89.23	94.06	95.56	---	103.04	107.02
13	101.83	101.00	---	---	88.84	87.30	89.39	93.92	95.71	---	103.03	106.89
14	102.05	101.14	---	88.48	88.85	---	89.95	93.92	95.91	---	103.07	106.79
15	102.10	100.83	---	---	88.87	87.45	90.36	93.95	96.17	---	103.17	106.98
16	102.35	100.61	---	---	88.98	87.44	90.31	94.02	96.45	---	103.33	106.94
17	102.57	99.19	---	---	89.19	---	90.28	94.15	96.68	---	103.49	106.87
18	102.68	96.19	---	---	89.49	---	90.27	94.28	96.87	---	103.59	106.84
19	102.66	94.01	---	---	89.57	---	90.30	94.38	97.10	---	103.67	106.90
20	102.92	92.47	---	---	89.34	---	90.41	94.38	97.43	---	103.88	107.01
21	103.07	91.56	---	---	89.30	---	90.60	94.40	98.19	---	104.00	107.16
22	103.26	89.80	---	---	89.30	86.86	90.79	94.47	98.49	---	104.22	107.31
23	103.28	84.60	---	---	89.27	87.01	90.99	94.58	98.48	---	104.78	107.38
24	102.32	81.94	---	---	89.37	87.31	91.11	94.78	98.67	---	105.14	107.45
25	99.48	---	---	89.69	89.50	87.57	91.09	95.04	---	---	105.09	107.59
26	97.45	---	---	90.12	89.57	87.51	91.24	95.31	---	---	105.08	107.76
27	97.00	---	---	89.97	89.53	87.54	91.58	95.55	---	---	105.14	108.00
28	97.07	79.72	---	89.71	89.47	87.57	92.15	95.90	---	---	105.26	108.24
29	97.32	79.88	---	89.65	---	87.79	92.08	95.48	---	---	105.40	108.54
30	97.95	80.35	---	89.63	---	87.99	92.22	95.26	---	---	105.57	108.75
31	98.45	---	---	89.70	---	88.10	---	95.22	---	---	105.80	---
Mean	101.22	---	---	---	89.39	---	89.95	94.24	---	---	---	107.17
Maximum	105.75	---	---	---	89.96	---	92.22	95.90	---	---	---	108.75
Minimum	97.00	---	---	---	88.84	---	88.10	92.37	---	---	---	106.10

Appendix 1.1. Daily mean depth to water at well DX-68-22-913, Comal County, Texas, 2004-06—Continued.

Water year 2006 (Oct. 2005–Sept. 2006) daily mean values												
Day	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	108.88	107.83	---	---	110.85	112.18	114.66	117.79	120.01	123.79	125.97	128.54
2	109.00	107.79	---	---	110.83	112.28	114.58	117.82	120.16	123.90	126.07	128.59
3	109.11	107.85	---	---	110.87	112.37	114.57	117.80	120.27	124.01	126.17	128.64
4	109.10	108.12	---	---	110.92	112.52	114.64	117.81	120.36	124.07	126.27	128.69
5	108.39	108.14	---	109.49	110.94	112.75	114.71	117.81	120.49	123.98	126.37	128.73
6	107.78	108.11	---	109.57	111.03	112.88	114.79	117.71	120.64	123.80	126.48	128.71
7	107.87	108.09	---	109.74	111.07	113.01	114.94	117.50	120.77	123.71	126.61	128.70
8	107.86	108.12	---	109.77	111.13	113.10	115.07	117.35	120.92	123.62	126.75	128.66
9	107.79	108.19	---	---	111.25	113.20	115.15	117.28	121.08	123.55	126.85	128.59
10	107.74	108.28	---	---	111.38	113.32	115.22	117.15	121.25	123.55	126.93	128.49
11	107.76	108.33	---	109.97	111.49	113.42	115.35	117.09	121.39	123.57	127.05	128.39
12	107.74	108.34	---	109.99	111.50	113.50	115.47	117.04	121.58	123.66	127.09	128.34
13	107.64	108.35	---	110.08	111.49	113.61	115.60	117.03	121.79	123.74	127.15	128.18
14	107.53	108.34	---	110.17	111.54	113.76	115.76	117.05	121.99	123.85	127.23	128.06
15	107.42	108.37	---	110.20	111.66	113.91	115.85	117.11	122.21	123.98	127.29	127.95
16	107.34	108.45	---	110.26	111.76	114.04	115.97	117.17	122.40	124.10	127.36	127.78
17	107.32	108.51	---	110.42	111.87	114.17	116.11	117.30	122.56	124.22	127.44	127.59
18	107.30	108.73	---	---	111.94	114.28	116.33	117.41	122.62	124.38	127.53	127.39
19	107.32	108.71	---	---	111.92	114.37	116.58	117.55	122.65	124.55	127.63	127.22
20	107.39	108.67	---	---	111.89	114.43	116.81	117.72	122.54	124.74	127.72	127.04
21	107.49	108.68	---	---	111.88	114.48	116.95	117.91	122.48	124.88	127.82	126.85
22	107.54	---	---	---	111.90	114.54	117.02	118.11	122.58	125.02	127.92	126.68
23	107.57	---	---	---	---	114.63	117.07	118.33	122.64	125.13	128.03	126.53
24	107.62	109.07	---	---	112.08	114.68	117.13	118.54	122.74	125.27	128.14	126.38
25	107.62	109.08	---	110.68	112.09	114.70	117.23	118.79	122.83	125.35	128.21	126.28
26	107.66	---	---	110.72	112.10	114.68	117.33	119.04	122.96	125.44	128.28	126.20
27	107.74	---	---	110.79	112.07	114.67	117.41	119.25	123.13	125.52	128.35	126.10
28	107.83	108.99	---	110.79	112.10	114.67	117.49	119.45	123.31	125.61	128.41	126.02
29	107.89	108.96	---	110.74	---	114.64	117.57	119.61	123.47	125.70	128.46	125.94
30	107.88	---	---	110.79	---	114.64	117.68	119.74	123.64	125.78	128.48	125.88
31	107.85	---	---	110.84	---	114.70	---	119.86	---	125.87	128.50	---
Mean	107.84	---	---	---	---	113.81	116.03	117.97	121.92	124.46	127.37	127.57
Maximum	109.11	---	---	---	---	114.70	117.68	119.86	123.64	125.87	128.50	128.73
Minimum	107.30	---	---	---	---	112.18	114.57	117.03	120.01	123.55	125.97	125.88

Appendix 1.2. Daily mean depth to water at well DX-68-23-502, Comal County, Texas, 2004-06.

[In feet below land surface; ---, not collected or computed]

Water year 2004 (Oct. 2003–Sept. 2004) daily mean values												
Day	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	---	---	---	---	---	---	32.40	30.44	31.13	27.62	27.93	29.71
2	---	---	---	---	---	---	32.36	30.25	31.50	27.36	28.13	29.75
3	---	---	---	---	---	---	32.14	30.29	31.75	27.04	28.38	29.76
4	---	---	---	---	---	---	31.99	30.34	31.88	26.79	28.51	29.69
5	---	---	---	---	---	---	31.97	30.32	31.77	26.70	28.64	29.43
6	---	---	---	---	---	---	31.99	30.32	31.76	26.88	28.75	29.37
7	---	---	---	---	---	---	31.92	30.30	32.00	26.96	28.57	29.43
8	---	---	---	---	---	---	31.88	30.08	32.09	26.96	28.48	29.49
9	---	---	---	---	---	---	31.80	29.94	31.63	26.99	28.58	29.52
10	---	---	---	---	---	---	31.60	30.01	30.35	26.82	28.74	29.55
11	---	---	---	---	---	---	31.38	30.08	30.02	26.69	28.84	29.59
12	---	---	---	---	---	---	31.36	30.03	29.66	26.76	28.94	29.48
13	---	---	---	---	---	---	31.43	29.97	29.31	26.91	29.04	29.60
14	---	---	---	---	---	---	31.41	29.89	29.31	26.97	28.95	29.67
15	---	---	---	---	---	---	31.36	29.67	29.35	27.09	28.89	29.58
16	---	---	---	---	---	---	31.36	29.47	29.33	27.19	29.07	29.55
17	---	---	---	---	---	---	31.19	29.54	29.34	27.14	29.29	29.59
18	---	---	---	---	---	32.29	31.04	29.68	29.37	27.07	29.44	29.61
19	---	---	---	---	---	32.29	31.14	29.70	29.27	27.26	29.48	29.53
20	---	---	---	---	---	32.14	31.22	29.79	29.22	27.53	29.55	29.62
21	---	---	---	---	---	32.01	31.25	29.91	29.38	27.75	29.56	29.75
22	---	---	---	---	---	32.15	31.28	29.82	29.58	27.92	29.44	29.78
23	---	---	---	---	---	32.28	31.29	29.78	29.60	27.95	29.51	29.82
24	---	---	---	---	---	32.28	31.18	29.97	29.52	27.76	29.58	29.83
25	---	---	---	---	---	32.27	30.96	30.18	29.46	27.67	29.57	29.76
26	---	---	---	---	---	32.28	30.84	30.34	29.25	27.61	29.60	29.69
27	---	---	---	---	---	32.12	30.85	30.50	29.03	27.80	29.65	29.79
28	---	---	---	---	---	32.00	30.78	30.67	28.93	27.92	29.58	29.90
29	---	---	---	---	---	32.16	30.72	30.68	28.73	27.99	29.35	30.00
30	---	---	---	---	---	32.29	30.62	30.68	28.10	28.04	29.50	30.02
31	---	---	---	---	---	32.37	---	30.84	---	27.97	29.66	---
Mean	---	---	---	---	---	---	31.42	30.11	30.05	27.33	29.07	29.66
Maximum	---	---	---	---	---	---	32.40	30.84	32.09	28.04	29.66	30.02
Minimum	---	---	---	---	---	---	30.62	29.47	28.10	26.69	27.93	29.37

Appendix 1.2. Daily mean depth to water at well DX-68-23-502, Comal County, Texas, 2004-06—Continued.

Water year 2005 (Oct. 2004–Sept. 2005) daily mean values												
Day	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	30.04	26.48	17.83	19.59	21.61	21.43	20.70	23.46	24.96	28.97	29.80	31.46
2	29.71	26.35	17.81	19.65	21.57	21.38	20.61	23.63	24.86	29.10	29.93	31.52
3	29.11	26.22	17.83	19.72	21.59	21.40	20.57	23.78	24.81	29.08	30.07	31.46
4	28.98	26.15	17.67	19.85	21.65	21.42	20.74	23.93	24.78	29.21	30.25	31.33
5	28.93	26.09	17.55	20.02	21.46	21.24	20.87	24.06	24.68	29.54	30.26	31.31
6	28.85	25.95	17.69	20.09	21.32	20.92	21.02	24.20	24.75	29.80	30.16	31.45
7	28.80	25.77	17.83	20.20	21.42	20.76	21.11	24.23	24.90	29.98	30.13	31.51
8	---	25.84	17.88	20.19	21.53	20.67	21.23	24.13	25.01	30.13	30.32	31.52
9	28.58	25.92	17.91	20.13	21.44	20.59	21.16	24.15	25.09	30.16	30.46	31.58
10	28.37	25.94	18.03	20.27	21.52	20.50	21.08	24.28	25.15	30.15	30.44	31.54
11	28.42	26.00	17.94	20.48	21.42	20.44	21.34	24.31	25.25	30.45	30.31	31.29
12	28.52	26.06	17.84	20.54	21.26	20.27	21.53	24.37	25.27	30.71	30.28	31.20
13	28.57	26.05	18.13	---	21.04	20.10	21.65	24.36	25.42	30.87	30.21	31.21
14	28.51	25.86	18.38	---	21.17	20.20	21.80	24.34	25.66	30.94	29.95	31.13
15	28.45	25.76	18.45	---	21.26	20.30	21.93	24.23	25.85	31.02	30.04	31.03
16	28.38	25.76	18.55	---	21.27	20.27	21.89	24.30	26.01	30.87	30.15	31.02
17	28.25	24.84	18.67	---	21.33	20.20	21.86	24.41	26.16	30.43	30.13	31.03
18	28.31	24.15	18.62	---	21.38	20.18	22.03	24.46	26.31	30.23	30.18	30.96
19	28.41	23.81	18.57	---	21.22	20.06	22.22	24.52	26.36	30.17	30.25	31.17
20	28.46	23.40	18.67	---	21.14	19.94	22.33	24.60	26.65	30.11	30.25	31.32
21	28.56	22.90	18.83	---	21.30	20.01	22.39	24.65	26.93	30.03	30.21	31.39
22	28.58	21.49	18.99	---	21.46	20.20	22.43	24.61	27.17	29.98	30.39	31.42
23	28.28	19.38	19.12	---	21.47	20.27	22.49	24.80	27.38	29.89	30.55	31.46
24	27.95	18.83	19.04	---	21.50	20.31	22.46	24.97	27.60	29.66	30.69	31.47
25	27.52	18.38	19.00	21.53	21.59	20.36	22.61	25.16	27.78	29.77	30.86	31.46
26	27.32	17.96	19.08	21.61	21.47	20.27	22.77	25.26	27.82	29.87	30.97	31.72
27	27.19	17.90	19.20	21.65	21.27	20.22	23.02	25.30	28.08	29.90	31.03	31.95
28	27.08	17.69	19.37	21.59	21.35	20.34	23.21	25.16	28.34	29.89	30.99	32.11
29	27.01	17.71	19.57	21.51	---	20.47	23.36	24.91	28.53	29.83	31.13	32.19
30	26.89	17.83	19.71	21.39	---	20.52	23.47	24.83	28.74	29.83	31.28	32.26
31	26.68	---	19.65	21.47	---	20.59	---	24.94	---	29.72	31.35	---
Mean	---	23.42	18.50	---	21.39	20.51	21.86	24.46	26.21	30.01	30.42	31.45
Maximum	---	26.48	19.71	---	21.65	21.43	23.47	25.30	28.74	31.02	31.35	32.26
Minimum	---	17.69	17.55	---	21.04	19.94	20.57	23.46	24.68	28.97	29.80	30.96

Appendix 1.2. Daily mean depth to water at well DX-68-23-502, Comal County, Texas, 2004-06—Continued.

Water year 2006 (Oct. 2005–Sept. 2006) daily mean values												
Day	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
1	32.28	31.04	32.73	32.62	33.55	34.87	36.45	38.99	41.46	45.04	46.20	47.72
2	32.26	31.09	32.79	32.71	33.59	34.97	36.38	---	41.42	44.66	46.32	47.64
3	32.38	31.20	32.78	32.88	33.71	35.11	36.58	39.19	41.41	44.54	46.47	47.31
4	---	31.28	32.78	33.04	33.80	35.21	36.69	39.23	41.39	44.41	46.66	47.30
5	33.48	31.23	32.87	33.17	33.80	35.24	36.84	39.07	41.63	43.74	46.61	47.42
6	31.33	31.20	32.79	33.25	33.86	35.49	36.92	38.67	41.97	43.51	46.34	47.03
7	30.42	31.42	32.78	33.23	33.96	35.62	37.02	38.20	42.18	43.38	46.32	46.82
8	30.30	31.56	32.71	33.12	34.03	35.73	37.11	38.03	42.40	43.19	46.45	46.72
9	30.28	31.64	32.80	33.25	34.09	35.78	37.08	37.97	---	43.11	46.53	46.39
10	30.34	31.72	32.74	33.28	34.13	---	37.29	37.97	---	43.31	46.64	46.06
11	30.35	31.77	32.62	33.26	34.16	---	37.50	38.04	---	43.62	46.83	46.03
12	30.23	31.77	32.67	33.35	34.06	---	37.70	38.08	---	43.94	46.82	45.85
13	30.17	31.73	32.71	33.47	34.20	---	37.85	38.18	---	44.23	46.61	45.51
14	30.16	31.91	32.73	33.41	34.36	---	38.01	38.19	---	44.48	46.84	45.32
15	30.11	32.02	32.78	33.44	34.46	36.80	38.19	38.35	---	44.47	47.10	45.19
16	30.05	32.11	32.72	33.49	34.49	36.79	38.21	38.59	---	44.46	47.28	44.97
17	30.21	32.17	32.65	33.54	34.61	36.88	38.43	38.90	---	44.78	47.41	44.72
18	30.28	32.20	32.53	33.54	34.66	36.89	38.71	39.19	---	45.24	47.55	44.54
19	30.35	32.18	32.52	33.50	34.55	36.83	38.94	39.41	---	45.47	47.43	44.39
20	30.41	32.10	32.60	---	34.55	36.82	39.15	39.62	---	45.61	47.30	44.27
21	30.51	32.22	32.63	33.54	34.60	36.83	39.08	39.74	43.68	45.70	47.41	44.15
22	30.56	32.40	32.67	33.38	34.61	36.80	38.83	40.12	43.78	45.68	47.60	44.12
23	30.52	32.50	32.51	33.42	34.67	36.81	38.74	40.47	43.96	45.41	47.60	44.01
24	30.71	32.43	32.44	33.60	34.73	36.80	38.92	40.76	44.00	---	47.64	43.73
25	30.76	32.37	32.39	33.66	34.71	36.71	39.09	41.04	43.97	45.64	47.72	43.73
26	30.84	32.40	32.43	33.61	34.60	36.55	39.17	41.29	44.19	45.71	47.62	43.84
27	30.90	32.34	32.56	33.65	34.68	36.59	39.26	41.53	44.53	45.84	47.32	43.84
28	31.00	32.50	32.70	33.60	34.81	36.57	39.23	41.46	44.76	45.95	47.39	43.85
29	31.02	32.64	32.78	33.45	---	36.48	39.13	41.25	44.99	45.92	47.63	43.89
30	30.86	32.69	32.71	33.43	---	36.45	38.91	41.23	45.22	45.75	47.52	43.85
31	30.98	---	32.63	33.53	---	36.48	---	41.39	---	45.94	47.57	---
Mean	---	31.93	32.67	---	34.29	---	38.05	---	---	---	47.06	45.34
Maximum	---	32.69	32.87	---	34.81	---	39.26	---	---	---	47.72	47.72
Minimum	---	31.04	32.39	---	33.55	---	36.38	---	---	---	46.20	43.73

Appendix 1.3. Chemical and isotope data in ground-water samples from wells and springs (by flow path) collected for this study, northeastern Bexar and southern Comal Counties, Texas, 2003–06.

[USGS, U.S. Geological Survey; °C, degrees Celsius; $\mu\text{S}/\text{cm}$, microsiemens per centimeter at 25 °C; mg/L, milligrams per liter; W, well; --, not analyzed for or not detected; dup, duplicate; Sp, spring; R, Rainfall; CaCO_3 , calcium carbonate; <, less than; E, estimated; $\mu\text{g}/\text{L}$, micrograms per liter; pptv, parts per trillion by volume; δD , delta deuterium; per mil, parts per thousand; $\delta^{18}\text{O}$, delta oxygen-18; ^3H , tritium; TU, tritium units; ^3He , helium-3; M, presence of material verified but not quantified]

Site type	State well number (USGS name)	Date	Time	Temperature, water (°C)	pH (standard units)	Specific conductance ($\mu\text{S}/\text{cm}$)	Dissolved oxygen (mg/L)	Calcium (mg/L)	Magnesium (mg/L)
Southern Comal flow path									
W	AY-68-29-610	7/28/2003	1200	22.5	7	608	--	100	13.5
		7/28/2003 dup	1201	--	--	--	--	--	--
W	AY-68-29-913	7/28/2003	1100	24.5	7.1	515	--	73.2	15.9
		7/28/2003 dup	1101	--	--	--	--	--	--
W	AY-68-29-914	7/28/2003	1000	25	7.2	493	--	68.2	16.1
		7/28/2003 dup	1001	--	--	--	--	--	--
W	DX-68-23-304	6/18/2003	1400	24	7.2	558	2.5	79.9	17
	(NBU-LCRA)	6/18/2003 dup	1401	--	--	--	--	--	--
		3/1/2006	1330	24.5	7.1	533	4.5	84.6	16.5
Sp	DX-68-23-324	7/22/2003	1230	23.9	7.1	547	--	84.7	17
	(Comal 7)	7/22/2003 dup	1231	--	--	--	--	--	--
		3/8/2006	1430	23.5	7.1	561	5.4	77.7	16.4
Sp	DX-68-23-325	8/7/2003	1030	--	--	--	--	79.4	17.8
	(Comal-Spring Island)	8/7/2003 dup	1100	--	--	--	--	98.7	18.5
		3/8/2006	1400	23.5	7.1	562	5.1	82.5	15.6
Sp	DX-68-23-326	7/22/2003	1330	24	7.1	548	--	86.8	17.6
	(Comal 5)	7/22/2003 dup	1331	--	--	--	--	--	--
		3/8/2006	1330	23.5	7.1	565	4.9	77.5	16.1
W	DX-68-23-602	7/24/2003	1100	23	7	544	--	83.2	13.5
		7/24/2003 dup	1101	--	--	--	--	--	--
Central Comal flow path									
W	AY-68-29-811	7/28/2003	0900	23.5	7	555	--	82.9	17.4
		7/28/2003 dup	0901	--	--	--	--	--	--
W	AY-68-30-520	7/17/2003	1100	26	7	527	5.5	74.9	16.7
		7/17/2003 dup	1101	--	--	--	--	--	--
Sp	DX-68-15-901	3/5/2003	1400	20	6.7	613	6.7	109	9.81
	(Hueco A)	3/5/2003 dup	1401	--	--	--	--	--	--
		3/16/2006	1300	20.5	7.1	594	5.5	86	17.9
Sp	DX-68-15-913	3/6/2003	1100	20	6.6	615	6.5	107	9.71
	(Hueco B)	3/6/2003 dup	1101	--	--	--	--	--	--

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Appendix 1.3. Chemical and isotope data in ground-water samples from wells and springs (by flow path) collected for this study, northeastern Bexar and southern Comal Counties, Texas, 2003–06—Continued.

Site type	State well number (USGS name)	Date	Time	Temperature, water (°C)	pH (standard units)	Specific conductance (µS/cm)	Dissolved oxygen (mg/L)	Calcium (mg/L)	Magnesium (mg/L)
Central Comal flow path—Continued									
W	DX-68-22-810	7/11/2003	1000	23	6.8	535	6.7	88.3	11.2
		7/11/2003 dup	1001	--	--	--	--	--	--
		3/14/2006	1300	22.8	7.1	544	5.8	98.6	6.56
W	DX-68-22-905	7/15/2003	1100	22.5	6.9	476	6.9	81	10.5
		7/15/2003 dup	1101	--	--	--	--	--	--
		3/16/2006	1100	21.5	7.2	534	6.3	85.8	11
Sp	DX-68-23-301 (Comal 1)	7/22/2003	1030	23.5	7	543	--	85.3	16
		7/22/2003 dup	1031	--	--	--	--	--	--
		3/1/2006	1110	23	7.2	558	5.3	80.9	17.1
Sp	DX-68-23-323 (Comal 3)	7/22/2003	1130	23.5	7	544	--	87.1	16.5
		7/22/2003 dup	1131	--	--	--	--	--	--
		3/1/2006	1150	23	7.2	556	5.2	81.9	16.2
W	DX-68-23-501	7/24/2003	1200	23.5	7	563	--	83.9	14.2
		7/24/2003 dup	1201	--	--	--	--	--	--
W	DX-68-23-710	7/15/2003	1200	23	6.7	588	7.5	110	8.75
		7/15/2003 dup	1201	--	--	--	--	--	--
		3/14/2006	1200	22.3	7.1	591	6.4	109	7.99
		3/14/2006 dup	1230	22.5	7.1	591	6.4	109	7.53
W	DX-68-30-111	7/17/2003	1000	22	6.7	664	6.3	115	8.2
		7/17/2003 dup	1001	--	--	--	--	--	--
Northern Comal flow path									
W	DX-68-23-104	7/9/2003	1300	23.5	7	527	4.3	62.4	32.6
		7/9/2003 dup	1301	--	--	--	--	--	--
Rainfall sites									
R	Rainfall site 3	6/5/2003	900	--	--	--	--	--	--
		7/8/2003	1400	--	--	--	--	--	--
R	Rainfall site 4	6/5/2003	900	--	--	--	--	--	--

Appendix 1.3. Chemical and isotope data in ground-water samples from wells and springs (by flow path) collected for this study, northeastern Bexar and southern Comal Counties, Texas, 2003–06—Continued.

State well number (USGS name)	Nickel (µg/L)	Silver (µg/L)	Zinc (µg/L)	Uranium (µg/L)	Sulfur hexa- fluoride (pptv)	δD (per mil)	δ ¹⁸ O (per mil)	³ H (TU)	³ He, tritogenic (TU)
Southern Comal flow path									
AY-68-29-610	3.01	<0.2	3.2	0.8	4.08	-23.09	-4.25	2.45	0
	--	--	--	--	3.98	--	--	--	--
AY-68-29-913	2.07	<.2	2.3	.75	3.9	-22.83	-4.03	2.23	-23.53
	--	--	--	--	3.52	--	--	--	--
AY-68-29-914	2.01	<.2	2.8	.69	1.57	-24.03	-4.22	1.99	-32.14
	--	--	--	--	1.77	--	--	--	--
DX-68-23-304 (NBU-LCRA)	1.47	<.2	6.8	.84	5.08	-24.14	-4.09	1.63	-57.05
	--	--	--	--	4.13	--	--	--	--
	1.68	<.2	1.2	.78	--	-23.83	-4.14	1.22	5.04
DX-68-23-324 (Comal 7)	2.59	<.2	E.7	.77	5.36	-23.31	-4.08	1.85	-38.23
	--	--	--	--	5.49	--	--	--	--
	1.53	<.2	4.5	.77	--	-22.67	-4.12	1.35	12.24
DX-68-23-325 (Comal-Spring Island)	2.47	<.2	<1.0	.82	4.29	-23.02	-4.23	2.13	-26.43
	2.46	<.2	<1.0	.81	4.18	--	--	--	--
	1.52	<.2	2	.74	--	-24.27	-4.13	1.26	5.67
DX-68-23-326 (Comal 5)	2.64	<.2	E.5	.78	4.66	-23.27	-4.05	2.03	-61.35
	--	--	--	--	4.51	--	--	--	--
	1.6	<.2	3.9	.76	--	-23.52	-4.15	1.33	7.57
DX-68-23-602	2.1	<.2	4	.68	2.98	-23.16	-4.14	2.21	.40
	--	--	--	--	3.82	--	--	--	--
Central Comal flow path									
AY-68-29-811	2.35	<.2	3	.81	15.69	-20.89	-3.92	2.13	-3.38
	--	--	--	--	7.69	--	--	--	--
AY-68-30-520	2.34	<.2	1.6	.78	6.33	-23.2	-4.12	1.93	-28.98
	--	--	--	--	6.31	--	--	--	--
DX-68-15-901 (Hueco A)	5.22	<.2	2.1	.88	3.7	-25.87	-4.54	1.98	-6.06
	--	--	--	--	3.47	--	--	--	--
	2.83	<.2	2.7	.93	--	-22.95	-3.96	1.53	-3.93
DX-68-15-913 (Hueco B)	4.89	<.2	3.1	.88	3.81	-24.46	-4.46	2.14	-6.03
	--	--	--	--	3.87	--	--	--	--

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Appendix 1.3. Chemical and isotope data in ground-water samples from wells and springs (by flow path) collected for this study, northeastern Bexar and southern Comal Counties, Texas, 2003–06—Continued.

State well number (USGS name)	Nickel (µg/L)	Silver (µg/L)	Zinc (µg/L)	Uranium (µg/L)	Sulfur hexa- fluoride (pptv)	δD (per mil)	δ ¹⁸ O (per mil)	³ H (TU)	³ He, tritogenic (TU)
Central Comal flow path—Continued									
DX-68-22-810	2.92	<0.2	642	0.65	1.96	-23.82	-4.28	2.30	-2.67
	--	--	--	--	1.77	--	--	--	--
	.28	<.2	796	.6	--	-23.2	-4.29	1.42	-2.64
DX-68-22-905	1.78	<.2	1.7	.63	2.88	-24.26	-4.18	2.21	.92
	--	--	--	--	2.75	--	--	--	--
	2.77	<.2	4.3	.69	--	-22.83	-4.33	1.60	-.08
DX-68-23-301 (Comal 1)	2.79	<.2	E.6	.74	4.33	-24.22	-4.13	2.03	-23.80
	--	--	--	--	4.32	--	--	--	--
	3.47	<.2	20.6	.84	--	-24.12	-4.16	1.25	4.84
DX-68-23-323 (Comal 3)	2.68	<.2	<1.0	.76	4.19	-23.36	-4.2	1.12	-49.90
	--	--	--	--	4.19	--	--	--	--
	1.42	<.2	1.5	.78	--	-23.22	-4.13	1.23	5.30
DX-68-23-501	2.07	<.2	10.6	.72	6.08	-22.66	-4.14	2.25	-14.22
	--	--	--	--	4.45	--	--	--	--
DX-68-23-710	2.11	<.2	1.7	.71	2.42	-25.3	-4.5	.34	.73
	--	--	--	--	2.07	--	--	--	--
	3.95	<.2	5	.61	--	-25.14	-4.58	1.23	.02
	2.98	<.2	4.9	.61	--	--	--	--	--
DX-68-30-111	4.4	<.2	375	.78	4.64	-24.63	-4.2	2.15	-.01
	--	--	--	--	5.16	--	--	--	--
Northern Comal flow path									
DX-68-23-104	1.51	<.2	21.4	1.02	1.16	-22.44	-3.86	.00	-15.28
	--	--	--	--	1.14	--	--	--	--
Rainfall sites									
Rainfall site 3	--	--	--	--	--	-35.76	-4.92	--	--
	--	--	--	--	--	-38.04	-6.02	--	--
Rainfall site 4	--	--	--	--	--	-41.7	-6.1	--	--

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