

Prepared in cooperation with the Arizona Department of Water Resources

# Groundwater Budgets for Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona, 2007–08

Scientific Investigations Report 2011–5159

U.S. Department of the Interior  
U.S. Geological Survey

**COVER:**

View of August 2010, monsoon thunderstorm over Hualapai Mountains, looking east from Sacramento Valley. Photo by Bradley D. Garner.

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By Bradley D. Garner and Margot Truini

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**U.S. Department of the Interior  
U.S. Geological Survey**

**U.S. Department of the Interior**  
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## Conversion Factors

<b>Inch/Pound to SI</b>		
<b>Multiply</b>	<b>By</b>	<b>To obtain</b>
<b>Length</b>		
inch (in.)	2.54	centimeter (cm)
inch (in.)	25.4	millimeter (mm)
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
<b>Area</b>		
square mile (mi <sup>2</sup> )	259.0	hectare (ha)
square mile (mi <sup>2</sup> )	2.590	square kilometer (km <sup>2</sup> )
<b>Volume</b>		
gallon (gal)	3.785	liter (L)
gallon (gal)	0.003785	cubic meter (m <sup>3</sup> )
gallon (gal)	3.785	cubic decimeter (dm <sup>3</sup> )
million gallons (Mgal)	3,785	cubic meter (m <sup>3</sup> )
cubic foot (ft <sup>3</sup> )	28.32	cubic decimeter (dm <sup>3</sup> )
cubic foot (ft <sup>3</sup> )	0.02832	cubic meter (m <sup>3</sup> )
acre-foot (acre-ft)	1,233	cubic meter (m <sup>3</sup> )
acre-foot (acre-ft)	0.001233	cubic hectometer (hm <sup>3</sup> )
<b>Flow rate</b>		
acre-foot per year (acre-ft/yr)	1,233	cubic meter per year (m <sup>3</sup> /yr)
acre-foot per year (acre-ft/yr)	0.001233	cubic hectometer per year (hm <sup>3</sup> /yr)
cubic foot per second (ft <sup>3</sup> /s)	0.02832	cubic meter per second (m <sup>3</sup> /s)
gallon per minute (gal/min)	0.06309	liter per second (L/s)
gallon per day (gal/d)	0.003785	cubic meter per day (m <sup>3</sup> /d)
million gallons per day (Mgal/d)	0.04381	cubic meter per second (m <sup>3</sup> /s)
inch per year (in/yr)	25.4	millimeter per year (mm/yr)

- Temperature in degrees Fahrenheit (°F) may be converted to degrees Celsius (°C) as follows: °C=(°F-32)/1.8
- Vertical coordinate information is referenced to the insert datum name (and abbreviation) here for instance, “North American Vertical Datum of 1988 (NAVD 88).”
- Horizontal coordinate information is referenced to the insert datum name (and abbreviation) here for instance, “North American Datum of 1983 (NAD 83).”
- Altitude, as used in this report, refers to distance above the vertical datum.

# Groundwater Budgets for Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona, 2007–08

By Bradley D. Garner and Margot Truini

## Abstract

The United States Geological Survey, in cooperation with the Arizona Department of Water Resources, initiated an investigation of the hydrogeology and water resources of Detrital, Hualapai, and Sacramento Valleys in northwestern Arizona in 2005, and this report is part of that investigation. Water budgets were developed for Detrital, Hualapai, and Sacramento Valleys to provide a generalized understanding of the groundwater systems in this rural area that has shown some evidence of human-induced water-level declines. The valleys are within the Basin and Range physiographic province and consist of thick sequences of permeable alluvial sediment deposited into basins bounded by relatively less permeable igneous and metamorphic rocks. Long-term natural recharge rates (1940–2008) for the alluvial aquifers were estimated to be 1,400 acre-feet per year (acre-ft/yr) for Detrital Valley, 5,700 acre-ft/yr for Hualapai Valley, and 6,000 acre-ft/yr for Sacramento Valley. Natural discharge rates were assumed to be equal to natural recharge rates, on the basis of the assumption that all groundwater withdrawals to date have obtained water from groundwater storage. Groundwater withdrawals (2007–08) for the alluvial aquifers were less than 300 acre-ft/yr for Detrital Valley, about 9,800 acre-ft/yr for Hualapai Valley, and about 4,500 acre-ft/yr for Sacramento Valley. Incidental recharge from leaking water-supply pipes, septic systems, and wastewater-treatment plants accounted for about 35 percent of total recharge (2007–08) across the study area. Natural recharge and discharge values in this study were 24–50 percent higher than values in most previously published studies. Water budgets present a spatially and temporally “lumped” view of water resources and incorporate many sources of uncertainty in this study area where only limited data presently are available.

## Introduction

In northwestern Arizona, residents of Detrital, Hualapai, and Sacramento Valleys are contending with a growing population and limited water resources in an arid region. As in many rural areas throughout the American desert southwest, surface-water resources are extremely limited (Jacobs and Stitzer, 2007), and groundwater is relied upon heavily to meet water

demand. Effects of human groundwater use are evident in parts of Detrital, Hualapai, and Sacramento Valleys, with water-level declines of more than 130 feet (ft) associated with groundwater withdrawals (Anning and others, 2007). An increasing population (about 30 percent from 2000 to 2007) most likely will place further burdens on groundwater resources. Recently proposed developments for approximately 200,000 new houses and condominiums (K. Davidson, Mohave County Planning and Zoning Department, written commun., 2008) and new solar-power plants (Arizona Department of Water Resources, 2009b, p. 55; Hawkins, 2009) would be additional potential stresses on the groundwater systems.

The U.S. Geological Survey (USGS), in cooperation with the Arizona Department of Water Resources (ADWR), initiated a hydrologic investigation of Detrital, Hualapai, and Sacramento Valleys in 2005 as part of the Rural Watershed Initiative (RWI; Anning and others, 2006), a program created by the State of Arizona and managed by ADWR to develop an improved understanding of rural water resources. As a part of this investigation, water budgets have been created for Detrital, Hualapai, and Sacramento Valleys and are presented in this report. Water budgets are used to answer fundamental hydrologic questions, such as how much groundwater exists, how much is entering an aquifer, and how much is exiting. Water budgets invariably are simplifications and contain uncertainty, but can be excellent tools for generalized, first-order understanding and management of complex hydrologic systems (Healy and others, 2007).

## Purpose and Scope

The purpose of this report is to present annual water budgets for each of Detrital, Hualapai, and Sacramento Valleys. Even small water-budget components with high uncertainty are identified and quantified where possible. The water budgets are “lumped” both spatially and temporally, meaning that single sets of numbers represent average combined effects for large areas and time spans. Although there is some discussion about spatial and temporal variability, the focus generally is on basin-wide conditions. The report concludes with a discussion of findings common to all three valleys, including uncertainty and limitations.

## 2 Groundwater Budgets for Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona, 2007–08

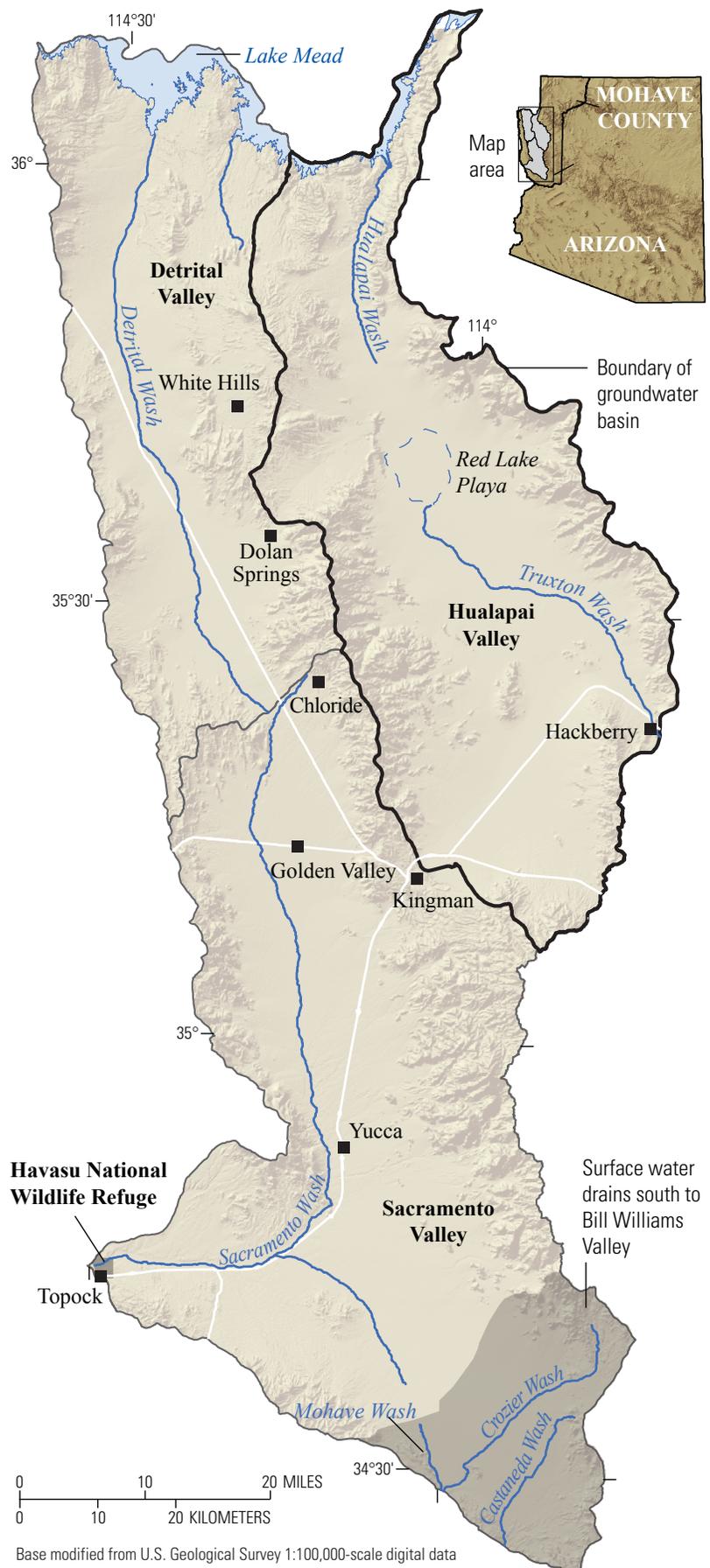
Water budgets in this report represent average conditions as of 2007–08, and therefore, they do not represent predevelopment or steady-state conditions. Groundwater withdrawals, incidental recharge, and changes in groundwater storage are calculated as average 2007–08 conditions. Natural recharge and discharge, however, are calculated as long-term averages for 1940–2008 because interannual variability of recharge is large in the study area, and the large volume of the alluvial aquifers can be thought of as averaging out this year-over-year variability.

Water budgets in this report are developed with respect to the alluvial groundwater aquifers in Detrital, Hualapai, and Sacramento Valleys. Surface and anthropogenic processes are considered only insofar as they aid in quantifying water entering and exiting the alluvial aquifers. Infrequent, storm-driven streamflow that can flow out of the valleys is not quantified and is beyond the scope of this report. As discussed throughout this report, a key assumption is that groundwater withdrawals to date have not yet had an effect on natural recharge or natural discharge rates.

### Description of the Study Area

The study area comprises Detrital, Hualapai, and Sacramento Valleys (fig. 1) within the Basin and Range physiographic province of North America (Fenneman, 1931). The region was shaped by Tertiary-period tectonism, wherein extensional faulting created horsts and grabens across a broad expanse of pre-existing (primarily metamorphic and igneous) rocks (Beard and others, in press; Fenneman, 1931). Thick sequences of alluvial fill, eroded from topographically high horst areas, were deposited into graben-formed basins beginning 13–16 million years ago (fig. 2). The sediment in these alluvial basins ranges in size from clays to boulders, with coarser grained material along basin margins and streambeds. Fine-grained material and evaporites generally are more prevalent in the basin centers. The thickness of alluvial fill is smallest near basin margins and reaches thicknesses of more than 14,000 ft in the center of Hualapai Valley (fig. 3). Permeability of alluvial fill varies by location and depth, but it generally is greatest at the basin margins and along stream channels and is smallest in basin centers.

Metamorphic and volcanic rock (hereafter, bedrock) define the margins and floors of the alluvial basins. The margins are topographically high



**Figure 1.** Physiography and location of Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona.

mountain ranges and hereafter are referred to as mountain blocks. Depth-to-bedrock analyses indicate that each valley within the study area consists of multiple hydrologically connected subbasins (Conway and Ivanich, 2007; Ivanich and Conway, 2007; Mason and others, 2007; fig. 3), each with distinctive hydrogeologic properties.

There are no major perennial stream reaches in the study area; the stream-channel network conveys water only briefly after large or intense rainfall or snowmelt events. Small springs in the mountain blocks (appendix 2) create short perennial-stream reaches, but ultimately any such spring discharge either returns to the atmosphere through the process of evapotranspiration (ET) or re-enters the groundwater system as recharge. Springs in downgradient areas of Detrital and Hualapai Valleys near Lake Mead discharge water that may be associated with bank-storage effects from Lake Mead (Laney, 1979).

Detrital Valley has an area of about 900 square miles ( $\text{mi}^2$ ), and its alluvial fill is as much as 5,000 ft thick; its largest ephemeral stream channel is Detrital Wash. Surface and groundwater flow is from south to north toward Lake Mead. Along its southern boundary, Detrital Valley is separated from Sacramento Valley by a topographic divide and shallow bedrock. There are three subbasins within Detrital Valley: Southern, Central, and Northern. As of the 2000 census, about 1,500 people lived in Detrital Valley, mostly in and around the community of Dolan Springs (U.S. Census Bureau, 2010a).

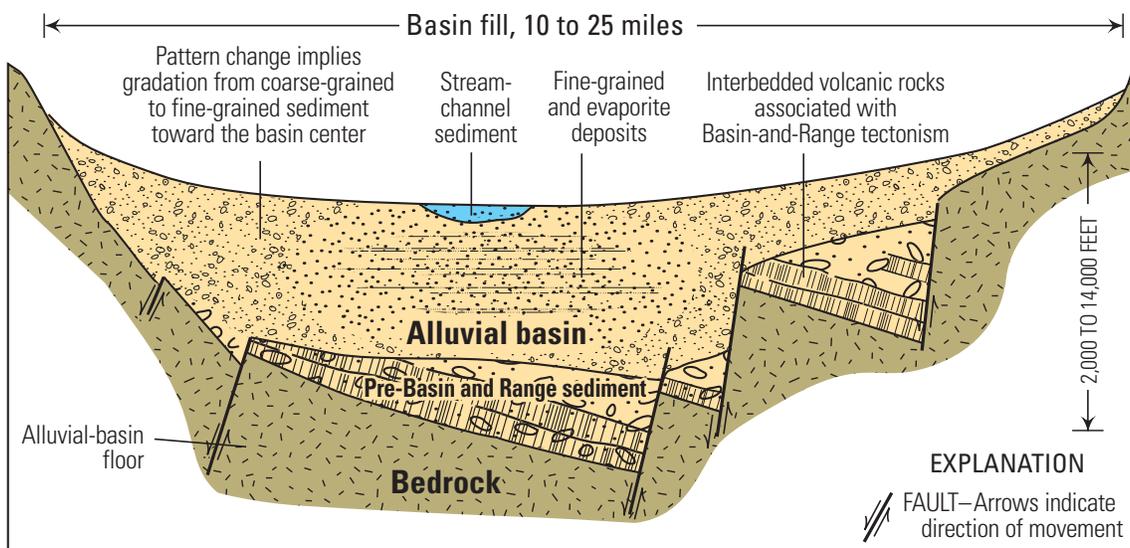
Hualapai Valley has an area of about 1,200  $\text{mi}^2$  and at its deepest point has almost twice the thickness of alluvial fill (more than 14,000 ft) as the other valleys in the study area. Truxton Wash is an ephemeral stream that enters Hualapai Valley from Truxton Canyon outside the study area and flows from southeast to northwest toward Red Lake Playa. This southern portion of Hualapai Valley is closed to surface-water drainage because of a topographic high north of Red Lake

Playa (Anderson, 1995). Groundwater in the valley flows generally from south to north toward Lake Mead. Red Lake is a vadose playa that is disconnected from the water table and, therefore, is not a groundwater discharge point (Anning and Konieczki, 2005). Hualapai Valley contains three subbasins: Kingman, Hualapai, and southern Gregg. Groundwater levels in some wells within the Kingman subbasin have declined more than 130 ft (Anning and others, 2007) because of City of Kingman groundwater withdrawals from this subbasin that began in 1979 (G. Sedich, City of Kingman, oral commun., 2010). As of the 2000 census, about 30,000 people lived in Hualapai Valley, mostly within and around the City of Kingman (U.S. Census Bureau, 2010a).

Sacramento Valley has an area of about 1,600  $\text{mi}^2$  and has about 8,500 ft of alluvial fill at its deepest point. Sacramento Wash, an ephemeral stream, flows from north to south and then west toward the Colorado River. Diversions and impoundments of the Colorado River near Havasu National Wildlife Refuge (NWR) have altered the groundwater-flow system near the downgradient end of Sacramento Valley (Guay, 2001). There are three subbasins in Sacramento Valley: Chloride, Golden Valley, and Dutch Flat. Water levels were declining in the Golden Valley subbasin from the 1950s until about 1979, when City of Kingman began to withdraw most of its water from the Hualapai subbasin of Hualapai Valley instead of the Golden Valley subbasin (see appendix 1). About 15,000 people lived in Sacramento Valley as of the 2000 census, mostly in the City of Kingman and the Golden Valley subdivision (U.S. Census Bureau, 2010a).

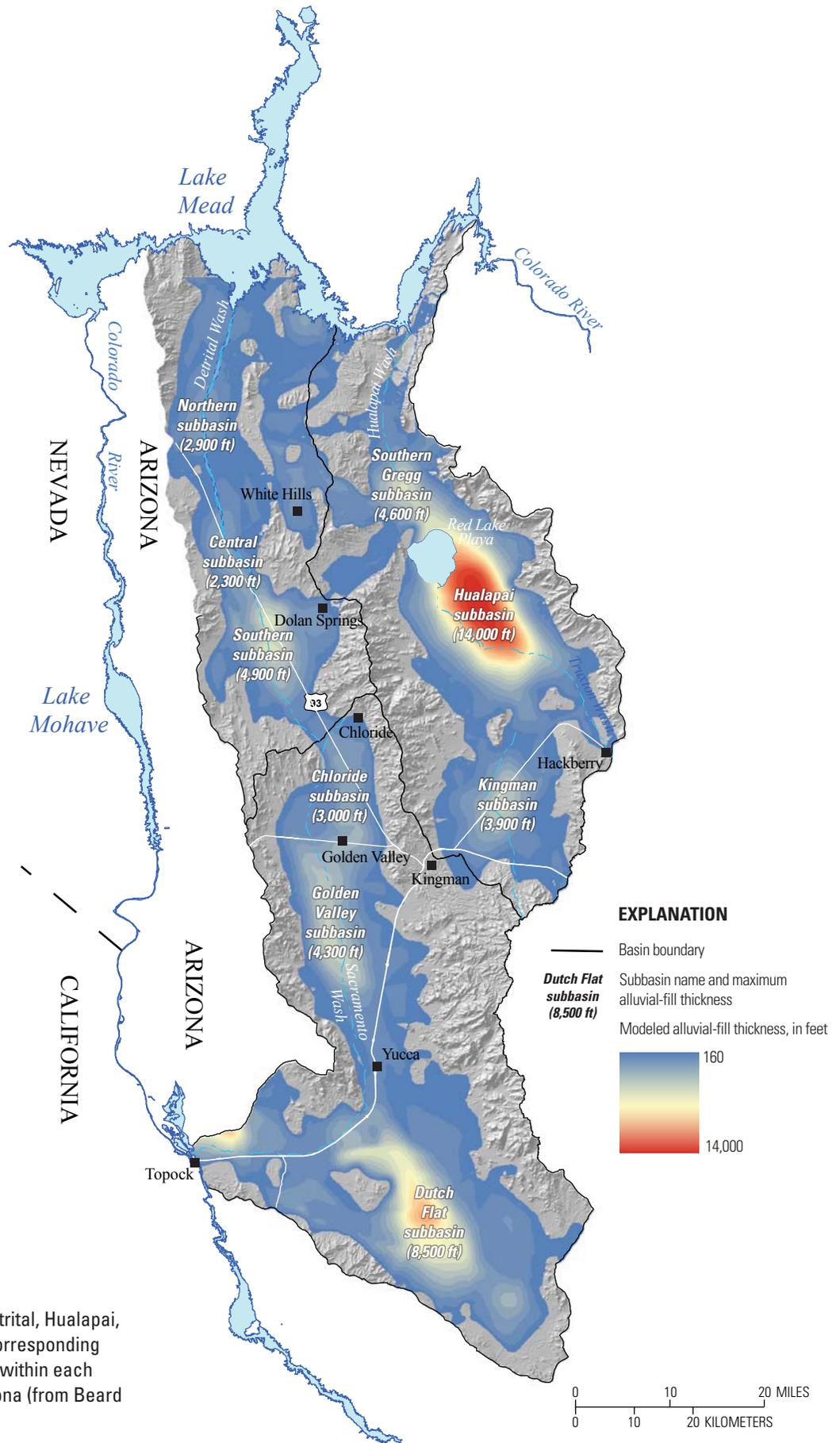
The study area has an arid to semiarid climate. Average precipitation is 9.2 inches per year (in/yr), on the basis of data from 1940 through 2008 (Johnson and Taylor, 2010; see “The Basin Characterization Model” section). Precipitation is higher over mountainous areas (10.2 in/yr) than over valley floors

## Mountain block



Vertical scale greatly exaggerated

**Figure 2.** Generalized schematic cross-section of the Basin and Range subsurface geology characteristic of Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona. Modified from Anderson (1995) and Beard and others (in press).



**Figure 3.** Subbasins within Detrital, Hualapai, and Sacramento Valleys, and corresponding maximum thickness of alluvium within each subbasin, Mohave County, Arizona (from Beard and others, in press).

(8.2 in/yr) because of elevation effects on precipitation (Daly and others, 2002). Precipitation occurs primarily in two seasons. Winter precipitation (normally November through February) is characterized by slow, steady precipitation lasting several days, as well as by occasional winter storms that generate large precipitation events. Summer precipitation (normally June through September) follows a monsoonal pattern, where brief, intense thunderstorms produce large amounts of precipitation (often 1 inch or greater) in short periods over small areas. Little precipitation falls between April and June, although meltwater from winter snow and ice in mountain areas can produce recharge and runoff during these months. Weather disturbances associated with tropical storms occasionally enter the study area in the summer and fall and can produce considerable precipitation.

Maximum daily air temperatures range from 50 to 70° F in the winter and from 90 to 110° F in the summer (Anning and others, 2007). High summertime temperatures greatly increase potential ET, consequently decreasing the amount of precipitation that can become groundwater recharge during these months.

## Previous Investigations

Two previous studies developed simplified water budgets for Detrital, Hualapai, and Sacramento Valleys. Gillespie and Bentley (1971) estimated surface-water runoff and groundwater-discharge rates for Hualapai and Sacramento Valleys. They also documented the occurrence of groundwater withdrawals and considered their impact on underflow between Hualapai Valley and Truxton Canyon outside of the study area. They did not evaluate incidental recharge. Freethy and Anderson (1986) presented predevelopment water budgets for all three valleys as part of the nationwide USGS Regional Aquifer-System Analysis (RASA) Program. Their values were considered to be semiquantitative only having been published using visual pie charts and arrows; however, the underlying quantitative values were available (AZ Water Science Center, U.S. Geological Survey, unpublished data) and, therefore, are shown for comparison in tables in this report.

Other studies contributed to an understanding of water-budget components in the study area, but those studies did not develop complete water budgets. Metzger and Loeltz (1973) estimated Sacramento Valley groundwater discharge as part of their study of the Needles, California, area. Rascona (1991) also estimated groundwater discharge from Sacramento Valley. Historical groundwater-withdrawal data have been published for all three valleys (Anning and Duet, 1994; Babcock and others, 1972; Babcock and others, 1970; Babcock and others, 1969; Cox and others, 1968; Gillespie and Bentley, 1971; Hodges and others, 1967; Pfaff and Clay, 1979; Tadayon, 2005; White and others, 1965; White and others, 1964; White and others, 1963) and are summarized in appendix 1. Estimates of total groundwater storage for all three valleys were calculated by Conway and Ivanich (2007), Ivanich and Conway (2007), Mason and others (2007), and a USGS study being undertaken for the RWI (M. Truini, U.S. Geological Survey, written commun., 2011).

## Acknowledgment

Don Bills, U.S. Geological Survey, provided assistance with data compilation and knowledge of the study area.

## Approach to Water-Budget Calculations

A water budget is based on conservation-of-mass accounting of water flowing into and out of a system (Healy and others, 2007):

$$I - O = \Delta S \quad (1)$$

where  $I$  is water input,  $O$  is water output, and  $\Delta S$  is change in water storage.

Although a water budget can be expressed entirely in equation form by expanding the terms in (1), lengthy and abstract equations can result. For this report, a reservoir-and-conveyance conceptual model was used that emphasized diagrammatic presentation (fig. 4). In this model, natural and human processes cause water to move between reservoirs along conveyances. An imbalance between total water conveyed into and out of a reservoir produces changes in storage in that reservoir. The reservoirs identified for this report were atmosphere, mountain block, ephemeral stream, alluvial unsaturated zone, alluvial aquifer, and human infrastructure. Processes identified that moved water between reservoirs were precipitation, ET, runoff, infiltration, recharge, upland spring discharge, groundwater withdrawals, consumptive use, water-supply pipe leakage, underflow, streamflow into the valley, and discharge to Lake Mead or the Colorado River.

Conceptually, any reservoir can undergo storage change, but for the purposes of this report only storage in the alluvial aquifer reservoir (hereafter, groundwater storage) was considered. Total storage and year-over-year storage change in snow, ice, perennial streams, lakes, vegetation, human infrastructure, and the alluvial unsaturated zone were beyond the scope of this report.

Water-budget component values are reported in units of acre-feet per year (acre-ft/yr), except for total groundwater storage, which is reported in millions of acre-feet (maf). Annual data are reported in terms of calendar years. Values greater than 10,000 acre-ft are rounded to the nearest 1,000 acre-ft, all others are rounded to the nearest 100 acre-ft, and values below 300 acre-ft are censored (less than, or <), unless reported more accurately by another source. Because of rounding and censoring, some table columns may not sum to zero.

Water budgets in this report represent average 2007–08 conditions, except for natural recharge, which was calculated for 1940–2008. The goal was to develop a single set of numbers that reflect recent (2007–08) human effects on alluvial aquifers, placed in the context of long-term (1940–2008) natural conditions. To that end, the following assumptions were made:

- **Natural recharge was unaffected by human activity during 1940–2008.** This was considered reasonable because

water tables in the study area are not near the land surface in recharge areas along mountain fronts and in alluvial streambeds (Anderson, 1995).

- **Withdrawal rates in 2007–08 were not affected by unusually high or low precipitation amounts.** This was considered reasonable, because withdrawals in 2007–08 are consistent with the overall increasing trend in withdrawals from 1940–2008 (see appendix 1).
- **Average climatic conditions were the same prior to 1940.** The period 1940–2008 includes multiple cycles of multiyear and decadal climate cycles known to exist in the southwestern United States (Pool, 2005). Longer scale climatic cycles and/or human-driven climate change would not necessarily be represented by this period, and are beyond the scope of this report.

### Natural Recharge (1940–2008)

Natural recharge to alluvial aquifers in the study area derives from precipitation that falls and infiltrates the ground and is not lost to ET or other processes. Also, some surface water and underflow enters Hualapai Valley from Truxton Canyon, a valley outside of the study area.

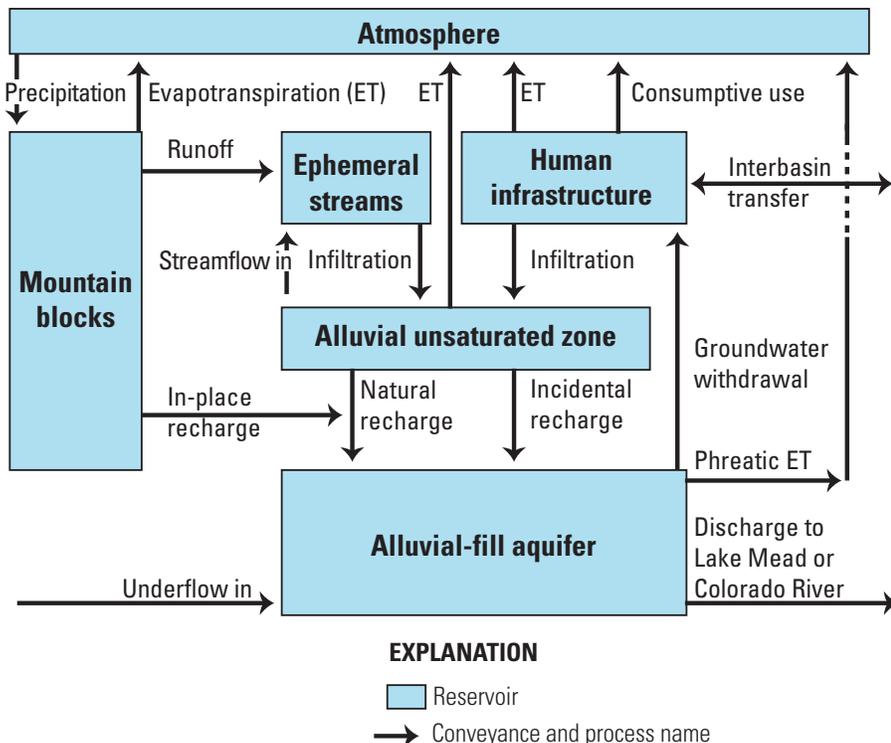
### The Basin Characterization Model

The Basin Characterization Model (BCM; Flint and Flint, 2007b) was used to estimate natural recharge to the alluvial

aquifers in the study area. The BCM is a physically based distributed model that calculates surface energy and surface mass flux in an 885-ft grid across the model domain. The model was developed for the entire southwestern United States; therefore, this report includes results only from a small portion of the model domain. Results from the model have been used in other hydrologic and water-balance studies (for example, Flint and Flint, 2007a; Heilweil and others, 2007; Tillman and others, 2011).

Natural recharge is a complex function of precipitation intensity, duration, magnitude, and temporal distribution, as well as geology, land-surface slope, and vegetative cover (Anderson, 1995). Inputs required for the BCM include topography, distributed precipitation and climate, and hydraulic properties of soil and surficial rocks. Precipitation and climate data were obtained from Parameter-elevation Regressions on Independent Slopes Model (PRISM) datasets from 1940–2008. PRISM datasets are model-derived, developed by extrapolating point measurements of precipitation at weather stations into a spatial grid by using statistical and human-expertise approaches (Daly and others, 2002; Johnson and Taylor, 2010). The number of precipitation stations available for use by PRISM was very limited within the study area compared to other areas of the United States, which increased uncertainty in precipitation estimates (Daly and others, 2008).

For each grid cell in the study area, a monthly volume of water was calculated and divided among three pathways: in-place recharge, runoff, and ET (fig. 5). In-place recharge is precipitation or meltwater that infiltrates and becomes groundwater recharge within the same grid cell in which it fell. Runoff is water remaining on the surface in a grid cell that could enter stream channels and eventually infiltrate through a channel streambed. Potential ET



**Figure 4.** Reservoir-and-conveyance conceptual model used to describe water budgets for Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona. Infiltration from the human-infrastructure reservoir includes septic tanks, wastewater-treatment plants, and leaking water-supply pipes. ET, evapotranspiration.

is subtracted from the amount of water calculated as runoff. BCM is not a routing model; it only calculates the water that infiltrates or remains on the surface in each grid cell. Grid cells can accumulate and store water as snowpack and soil water in one month and release this water in subsequent months.

The BCM indicates that, as in other arid alluvial basins, high potential ET within the study area generally prevents recharge and runoff from alluvial valley floors (Heilweil and others, 2007; Moore, 2007; Phillips, 1994; Stewart-Deaker and others, 2007; Stonestrom and others, 2007). Even though valley floors produce no recharge or runoff according to the BCM, runoff produced in mountain blocks flows ephemerally in stream channels that cross valley floors, and some of this runoff becomes recharge. In-place recharge from mountain blocks eventually either flows into alluvial aquifers through the subsurface, or reemerges from upland mountain-block springs.

### Calculation of Natural Recharge Using the Basin Characterization Model

Annual natural recharge from precipitation was estimated by using outputs of the BCM from 1940 to 2008 with the equation:

$$R = (R_i - Q_{us}) + 0.1 (RO + Q_{us}) \quad (2)$$

where  $R$  is natural recharge,  $R_i$  is BCM-computed in-place recharge,  $RO$  is BCM-computed runoff,  $Q_{us}$  is upland spring discharge, and 0.1 (10 percent) is a coefficient for how much runoff becomes recharge.

The assumption that 10 percent of runoff ( $RO$ ) becomes recharge was based on values from comparable areas summarized by Flint and Flint (2007b): 12–15 percent (Amargosa River, southern Nevada), 15–40 percent (Walnut Gulch, southeastern

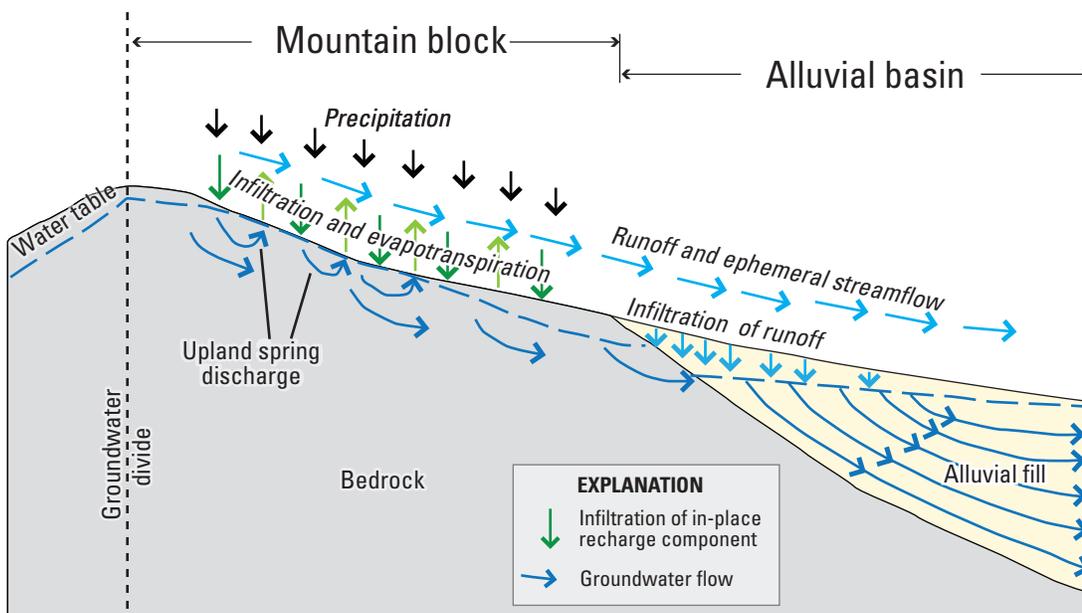
Arizona), 10 percent (southern Mojave Desert, California), 10 percent (Death Valley, California), and 90 percent (middle Humboldt Basin, western Nevada). The study area for this report is most similar to the study from the Mojave Desert of California.

Discharge from springs in upland mountain-block areas ( $Q_{us}$ ) was estimated by using an inventory of springs and spring-discharge measurements (appendix 2).  $Q_{us}$  was subtracted from  $R_i$  and added to  $RO$  under the assumption that upland springs represent a conversion of already accounted for in-place recharge into runoff.

### Natural Recharge from Upgradient Streamflow and Underflow

Groundwater and surface water in Hualapai Valley are connected hydrologically to an area outside of the study area known as Truxton Canyon. Average streamflow (runoff) entering Hualapai Valley in Truxton Wash was estimated by using data from USGS gaging station 09404343 (1994–2009). Of this streamflow, 45 percent was assumed to recharge the alluvial aquifer, a value halfway between 10 percent (BCM runoff) and 80 percent (incidental recharge). This intermediate value was arbitrarily chosen because alluvial streambeds are efficient at recharging aquifers because they are underlain by coarser-grained material (D. Pool, U.S. Geological Survey, oral commun., 2010), yet some of the infiltrating water likely returns to the atmosphere through ET within and adjacent to the streambed.

There is also a subsurface hydrologic connection—and, therefore, a potential for underflow—between Hualapai Valley and the adjacent Truxton Canyon, and this was evaluated by using groundwater levels, well logs, and previous studies.



**Figure 5.** Schematic diagram showing general groundwater-recharge processes in Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona. Modified from Berger (2000).

## Natural Discharge (1940–2008)

Natural discharge occurs from Detrital, Hualapai, and Sacramento Valley alluvial aquifers primarily as discharge to Lake Mead or the Colorado River (Freethey and Anderson, 1986; Gillespie and Bentley, 1971). Some natural discharge can occur as ET, which includes transpiration by vegetation that uses groundwater directly (phreatic ET). Other natural-discharge pathways either are not present in the study area (for example, perennial streamflow), or are assumed to be either negligible or irrelevant for the purposes of a basin-scale water budget (for example, downward and lateral leakage through bedrock).

## Discharge to Lake Mead or the Colorado River

Discharge to Lake Mead or the Colorado River is the primary means by which groundwater discharges from the alluvial aquifers within the study area. This component was assumed to be equal to natural recharge minus any phreatic ET, reflecting a key assumption discussed further in the “Effects of Groundwater Withdrawals” section.

Previous studies, particularly Freethey and Anderson (1986), have referred to this component as underflow. Although perhaps a reasonable term on a regional scale, underflow was considered a misnomer for this report; strictly speaking, the Colorado River is a perennial stream that gains groundwater discharge from the valleys.

Alternative approaches to calculating this discharge component were investigated, but ultimately not used. Springs along the perimeter of Lake Mead suggested a possible means of directly estimating natural discharge, but were not used for water-budget calculations (see the “Springs Near Lake Mead” section). Darcy’s law of groundwater flow was evaluated for applicability, but input parameters (hydraulic conductivity, gradient, and cross-sectional area) had too much uncertainty to produce a helpful, constraining result.

## Phreatic Evapotranspiration

Values for phreatic ET were obtained from Tillman and others (2011), a regional water-availability study of southwest alluvial basins that used greenness indexes from spectroradiometric satellite data to estimate phreatic ET along stream channels and in areas of woody and herbaceous plants. Tillman and others determined general water availability over an area much larger than Detrital, Hualapai, and Sacramento Valleys; therefore, their basin-scale results were evaluated for inclusion in the water budgets of this report using depth-to-water maps and field vegetation surveys.

## Groundwater Withdrawals

Groundwater wells throughout the study area are sources of water for domestic, municipal, industrial, and other water uses (fig. 6). More than 3,300 wells were registered with ADWR within the study area as of June 2009 (Arizona

Department of Water Resources, 2009a), but withdrawal data generally are not available on a per-well basis. Some organizations report total annual withdrawals of their actively pumped wells to State agencies, while the remaining wells generally are owned by individuals who are not required to report their withdrawals. It is unlikely that all 3,300 wells were in active use as of 2007–08; figure 6, therefore, is only a general indication of the location of withdrawals within the study area.

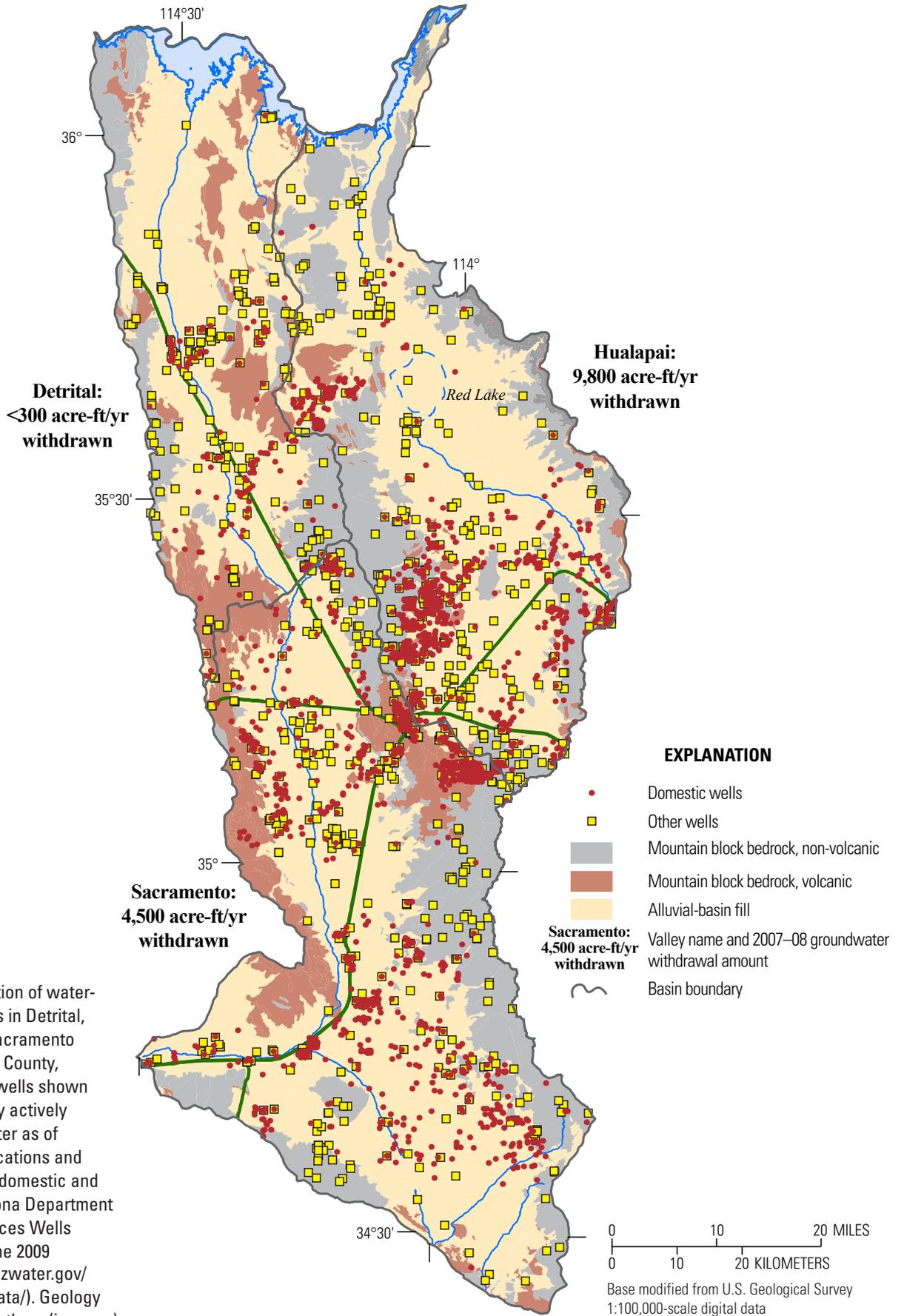
Annual withdrawal data (2007 and 2008) for the City of Kingman, community water suppliers (CWSs), and industrial facilities were compiled using methods described in Tadayon (2005). Data were obtained from reports submitted to the Arizona Corporation Commission (ACC; Arizona Corporation Commission, 2010), and questionable values were verified with telephone calls to water suppliers, where possible. These reported values were based on flowmeter measurements from 2007 and 2008, and values in this report are an average of these two years. Some water withdrawn by the City of Kingman from Hualapai Valley is transferred to Sacramento Valley to satisfy water demand (G. Sedich, City of Kingman, oral commun., 2010). The amount of water transferred was calculated by assuming a uniform rate of water use by City of Kingman residents, and dividing groundwater-withdrawal values between the valleys on the basis of City of Kingman population distribution between the two valleys (U.S. Census Bureau, 2010a).

Groundwater withdrawals by the self-supplied domestic (SSD) population were estimated indirectly by using population data, as these residents own small private wells with withdrawals that are neither measured nor reported. The SSD population was inferred by subtracting the number of people served by City of Kingman and CWSs in a valley from the estimated 2007 population of that valley. This inferred SSD population was multiplied by a per-capita water use of 133 gallons per person per day (Tadayon, 2005). The population served by City of Kingman municipal utility in each valley was assumed to be the census-reported “urban” population in that valley (no City of Kingman residents live in Detrital Valley). The population served by CWSs in each valley was estimated by dividing total annual CWS water deliveries by a per-capita water use of 133 gallons per person per day.

Study-area population estimates were obtained from Year 2000 U.S. Census data (U.S. Census Bureau, 2010a) that were adjusted to estimated 2007 levels by increasing the population 29 percent uniformly across the study area. This percentage was calculated on the basis of 2007 population estimates for City of Kingman as well as Mohave County (U.S. Census Bureau, 2010b).

## Incidental Recharge

Water use by humans can result in recharge when some of the nonconsumptively used water returns to the groundwater system. In a populated arid basin, this incidental recharge can be a substantial proportion of total recharge. Although no direct measurements of incidental recharge are available for the study area, potential pathways of incidental recharge (fig. 4) were investigated. Water is



**Figure 6.** Location of water-withdrawal wells in Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona. Not all wells shown were necessarily actively withdrawing water as of 2007–08. Well locations and water-use type (domestic and other) from Arizona Department of Water Resources Wells 55 database, June 2009 (<https://gisweb.azwater.gov/waterresourcedata/>). Geology from Beard and others (in press).

delivered to customers through water-supply pipes that invariably leak. Some water delivered to customers is returned to the atmosphere through consumptive use. The remaining unconsumed water (wastewater) discharges to septic or sewer systems, and some of it eventually reenters the groundwater system as recharge.

### Water-Supply Pipe Leakage

The annual percentage of water that leaked from water-supply pipes (unmetered loss) was calculated as the difference between volume of groundwater withdrawn and volume of water billed as being delivered to customers. Some activities, such as firefighting and filling of municipal pools, contribute to unmetered loss, but were considered negligible. For City of Kingman municipal utility, 2007–08 data indicate average annual unmetered losses of 7 percent. For CWSs, average annual unmetered losses were about 9 percent, on the basis of data reported to the ACC that were screened for common data errors. These percentages fall within ranges observed worldwide (5 to more than 60 percent; Sharp and Garcia-Fresca, 2003), and are similar to the 10-percent value observed in water systems that have efficient water-main repair and replacement procedures (Sharp and Garcia-Fresca, 2003). Industrial and SSD water-supply pipes were assumed to have no leaks as there were no data to indicate otherwise, and their typically short lengths might be more actively monitored and maintained than municipal water-supply pipes. Water transferred from Hualapai to Sacramento Valley was assumed to leak from pipes only in Sacramento Valley.

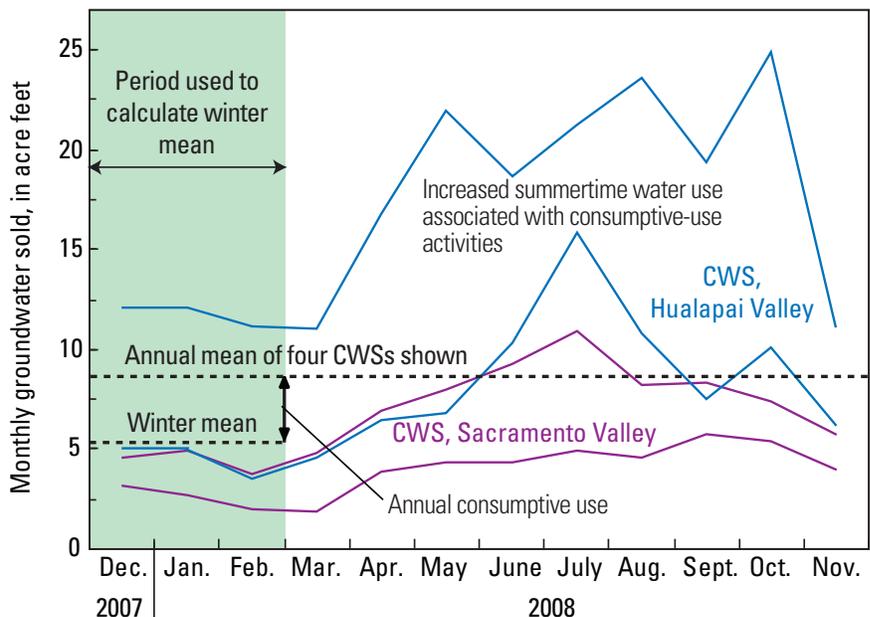
Of all unmetered-loss water, 80 percent was assumed to become recharge. The rationale for this assumption was that a leaking water pipe buried several feet below land surface behaves similarly to a septic-system drain field, and 80 percent was the amount of septic-system water assumed to become recharge. This is likely an oversimplification, but there have been no studies of recharge from leaking water-supply pipes in this or comparable areas.

### Consumptive Use

Some water delivered to customers for domestic use is returned to the atmosphere through consumptive use. Within the study area, consumptive-use activities consist primarily of summertime activities, such as yard watering, gardening, and use of evaporative-cooling systems. An annual consumptive-use coefficient of 28 percent was calculated for domestic water uses within the study area by using the winter base-rate method (LaTour, 1991; Shaffer, 2009). This method assumes that (1) all increased summertime domestic water use is associated with 100 percent consumptive-use activities, (2) consumptive use is minimal in the winter, and (3) consumptive use remains constant throughout the year. Monthly groundwater-withdrawal data were compiled from CWSs and screened for common data errors (for example, obvious hundred- or thousand-fold errors caused by incorrect reporting units). The coefficient was calculated as the mean winter-withdrawal rate for each year (December through February) subtracted from the annual-mean withdrawal rate for the entire year, divided by the annual-mean withdrawal rate (fig. 7). Consumptive-use coefficients calculated for 2007 and 2008 for 15 CWSs were between 18 and 43 percent, with a mean of 28 percent.

A consumptive-use coefficient of 100 percent was assumed for industrial water uses, indicating that all water supplied to industrial customers was expected to return to the atmosphere. This was considered reasonable because more than 90 percent of industrial water use in the study area was for thermoelectric power generation, a process that ultimately evaporates all water either through cooling towers or evaporation ponds (B. Henderson, Griffith Energy LLC, oral commun., 2010).

**Figure 7.** Monthly groundwater sales data reported by four community water suppliers (CWS) in Hualapai and Sacramento Valleys, Mohave County, Arizona. The figure is annotated to indicate how the winter base-rate method can be used to estimate consumptive use. Additional CWS data are available, but they are omitted here to improve readability.



## Recharge from Septic Systems

Most residents in the study area discharge their wastewater (that is, water not consumptively used) to septic systems, which in turn discharge to the unsaturated zone through drain fields. After a review of studies in comparable areas, 80 percent of this drain-field discharge was assumed to become recharge. Studies in Arizona’s San Pedro Valley used values of 100 percent (Arizona Department of Water Resources, 2005) and 60 percent (K. Lansey, University of Arizona, written commun., 2010). A study of California’s Joshua Tree groundwater basin assumed a value of 100 percent (Nishikawa and others, 2004), although the initial wetting front may take decades to reach the water table (Flint and Ellett, 2004).

The 20 percent of drain-field discharge that does not become recharge is accounted for in three ways. First, some initial amount of drain-field discharge is incorporated permanently into a very dry unsaturated zone. Second, some desert plants likely use some of this discharge water, even if it is discharged below the commonly assumed base of the root zone. Finally, upward gradients for water vapor can exist in arid unsaturated zones (Stonestrom and others, 2007; Walvoord and others, 2004), although fluxes through this path usually are much smaller than recharge values (de Vries and Simmers, 2002).

## Recharge from Wastewater-Treatment Plants

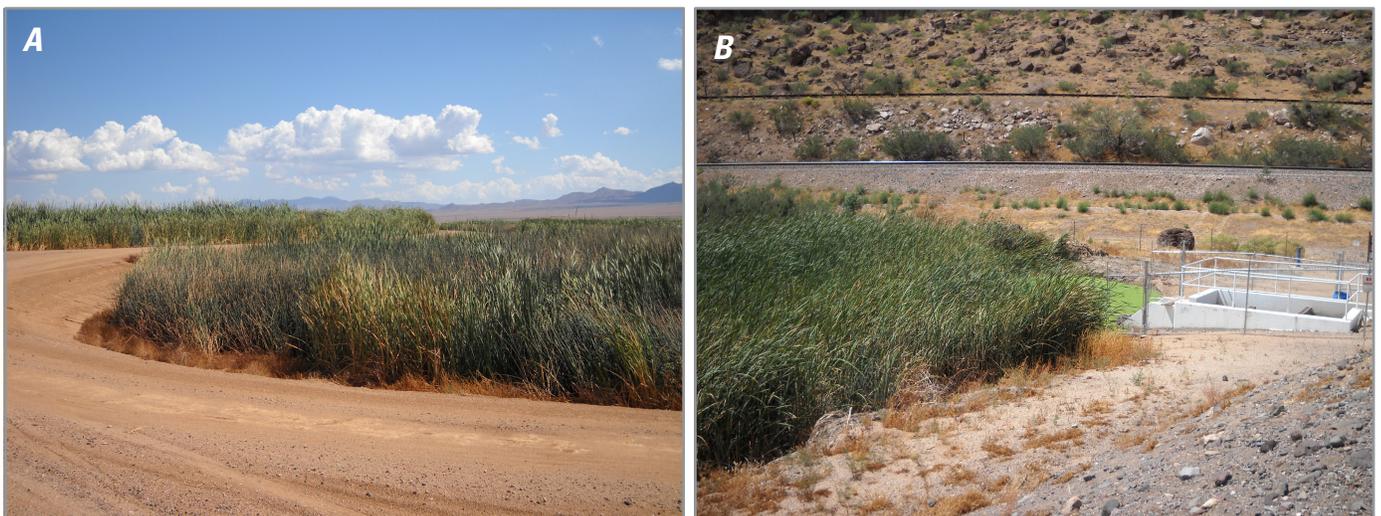
About 40 percent of City of Kingman residents discharge their wastewater to sewer systems that lead to two wastewater-treatment plants (WWTPs). The other 60 percent of City of Kingman residents discharge wastewater to septic systems; this proportion was determined by comparing the volume of unconsumed water calculated as being produced in City of Kingman with the volume of water reported as being treated by the WWTPs.

At City of Kingman’s two WWTPs, treated effluent is discharged into stands of dense vegetation (fig. 8). Treated effluent from the WWTP in Hualapai Valley is discharged to constructed wetlands, and the WWTP in Sacramento Valley discharges treated effluent to a stream channel filled with vegetation. One-hundred percent of treated effluent likely returns to the atmosphere through ET during June, July, and August as no surface-water flow exits these vegetation stands during these months (J. Corwin, City of Kingman, oral commun., 2010), and infiltration is unlikely because of clogging layers (Treese and others, 2009) or soil compaction. Fifty percent of effluent is assumed to return to the atmosphere through ET during April, May, September, and October; ET is assumed to be 10 percent of the treated effluent during all other months. After weighing each of these percentages (100, 50, and 10) by the proportional amount of water treated in each month, about 52 percent of treated wastewater was estimated to return to the atmosphere through ET on an annual basis, leaving the remaining 48 percent to infiltrate and potentially become recharge.

Of the 48 percent of treated effluent assumed to infiltrate annually, 80 percent is assumed to recharge the alluvial aquifer. This is the same percentage assumed for discharge from septic systems, and the rationales are similar. This water recharges either through streambed infiltration (Sacramento Valley WWTP), or through rapid infiltration basins (Hualapai Valley WWTP).

## Groundwater Storage and Storage Changes

Within the study area, most groundwater storage is within the intergranular pore spaces of the alluvial aquifers of each valley. Fractures within bedrock (particularly volcanic rocks) also store small but sometimes important volumes of water. The volume of water stored in bedrock is assumed to be negligible compared to the large amount stored in the alluvial



**Figure 8.** A, Cattails and bulrush growing in 75 acres of standing water in constructed wetlands at the City of City of Kingman wastewater-treatment plant in Hualapai Valley, Mohave County, Arizona; B, Discharge of about 300,000 gallons per day of treated effluent into Holy Moses Wash from the City of City of Kingman wastewater-treatment plant in Sacramento Valley, Mohave County, Arizona.

aquifers, and therefore, groundwater-storage estimates in this report do not include storage in bedrock.

Estimates of total groundwater storage for the alluvial aquifers of Detrital, Hualapai, and Sacramento Valleys to a depth of 1,300 below land surface (bls) have been calculated for a separate USGS study for the RWI (M. Truini, U. S. Geological Survey, written commun., 2011) and are shown in this report. They were based on 2006 water-table elevations and include any storage reductions that have been caused by withdrawals up through 2005.

Groundwater-storage changes in the alluvial aquifers of Detrital, Hualapai, and Sacramento Valleys were calculated by subtracting output water-budget components from input components. Inputs comprised natural recharge and incidental recharge. Outputs comprised natural discharge (phreatic ET and discharge to Lake Mead or the Colorado River) and groundwater withdrawals (fig. 4).

Because natural discharge is assumed to equal natural recharge, storage changes represent only the imbalance between groundwater withdrawals and incidental recharge. One implication of this assumption is that groundwater withdrawals are assumed to derive all water from depletion of groundwater storage and none by depletion of natural discharge or streamflow. This key assumption is discussed further in the “Effects of Groundwater Withdrawals” section.

A method for estimating storage change by using gravity measurements (Pool and Eychaner, 1995) was considered but not used, because there were not enough long-term data available from gravity-measurement stations in the study area. Changes between consecutive measurements in 2007 and 2008 were within measurement uncertainty (R. Carruth, U.S. Geological Survey, oral commun., 2010), except for one station in City of Kingman where a statistically significant decrease was consistent with water-level declines observed in that area.

## Detrital Valley Water Budget

### Natural Recharge

Natural recharge to the Detrital Valley alluvial aquifer is estimated to be 1,400 acre-ft/yr (fig. 9; table 1), on the basis of BCM calculations. Detrital Valley receives an average of 360,000 acre-ft/yr (7.6 in/yr) of precipitation. Of this total precipitation, 232,000 acre-ft/yr falls on the valley floor, and 128,000 acre-ft/yr falls on mountain blocks. All valley-floor precipitation returns to the atmosphere through ET, and about 1,400 acre-ft/yr (1.1 percent) of mountain-block precipitation eventually becomes recharge. This recharge comprises 1,200 acre-ft/yr of in-place recharge and 200 acre-ft/yr of recharge resulting from runoff.

### Natural Discharge

Natural discharge from the Detrital Valley alluvial aquifer is estimated to be 1,400 acre-ft/yr (fig. 9). This value was assumed to be the same as estimated natural recharge, reflecting a key assumption discussed further in the “Effects of Groundwater Withdrawals” section. Although most natural discharge occurs as groundwater discharge to Lake Mead, a small amount may occur as phreatic ET.

Tillman and others (2011) estimated that between 1,200 and 2,200 acre-ft/yr of groundwater discharges from the Detrital Valley alluvial aquifer through phreatic ET, but a spatial analysis of their results indicates that this range is unlikely to be accurate for Detrital Valley. Tillman and others estimated that most phreatic ET occurs along the shoreline of Lake Mead, but groundwater near the shoreline could well be bank-storage water, not water from the Detrital Valley alluvial aquifer. Tillman and others also estimated a small amount of phreatic ET occurring within and along the ephemeral Detrital Wash; however, groundwater levels from wells near Detrital Wash indicated a 32- to 870-ft unsaturated zone (Anning and others, 2007), suggesting that phreatic ET in this area is minimal. Thus, for the purposes of this water budget, phreatic ET was considered to be a negligible component of natural discharge.

Evidence of groundwater discharge or paleodischarge was found along the western side of Temple Bar Road in northern Detrital Valley, where small tufa mounds with distinctive vegetation appear to follow a linear strike (L.S. Beard, U.S. Geological Survey, oral commun., 2010). Tufa, a type of carbonate deposit, can indicate the discharge of mineralized groundwater. However, in the American southwest, tufa deposits most likely “indicate a change from a wetter to a drier climate” (Sweeting, 1973, p. 110), not a present-day groundwater discharge site. Distinctive vegetation on these mounds could be caused by differing soil chemistry and permeability rather than availability of shallow groundwater. Consequently, any present-day discharge at these mounds, considered unlikely in any case, is deemed negligible.

### Springs near Lake Mead

Springs were identified near the shoreline of Lake Mead that could be discharge points for groundwater from the Detrital Valley alluvial aquifer (see appendix 2). Because of the bank-storage effect of Lake Mead, however, care must be taken in interpreting this spring discharge and its possible sources. This section discusses the possibility of bank-storage water being the source of discharge for springs located below the maximum water-surface elevation of Lake Mead (1,229 ft above sea level).

As Lake Mead’s water-surface elevation rises, water infiltrates from the reservoir into permeable unsaturated sediments along the shoreline. When the water-surface elevation falls, water that had infiltrated into the sediments discharges back to the lake (Bales and Laney, 1992). This discharge can occur as diffuse seepage along canyon

**Table 1. Groundwater-budget values for Detrital Valley, Mohave County, Arizona; includes values from previous studies for comparison.**

[acre-ft/yr, acre-feet per year; maf, millions of acre-feet; <, less than; ET, evapotranspiration]

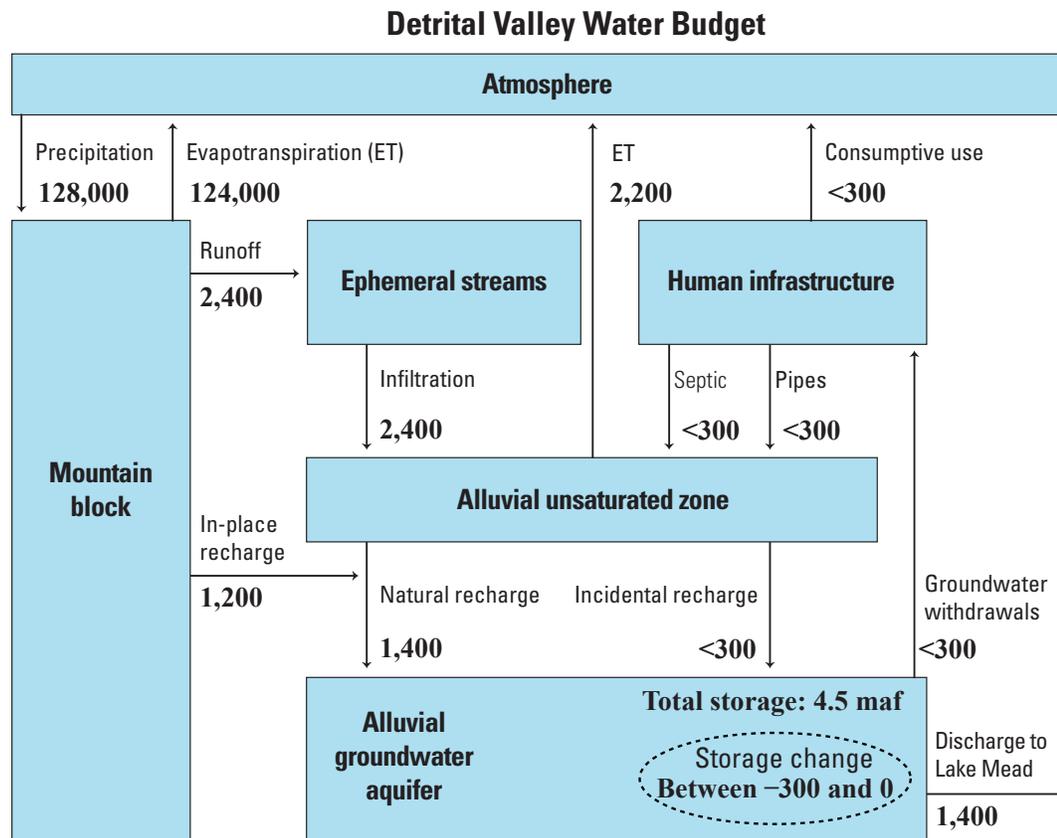
Water-budget component	This study		Other studies		
	Inflow to aquifer (acre-ft/yr)	Outflow from aquifer (acre-ft/yr)	Freethy and Anderson (1986) <sup>1</sup> (acre-ft/yr)	Tillman and others (2011) (acre-ft/yr)	Other (see footnote <sup>5</sup> ) (acre-ft/yr)
Natural recharge	1,400		700	<sup>2</sup> 2,000	
Mountain-block recharge	1,200				
Ephemeral stream-channel recharge	<300				
Natural discharge to Lake Mead		1,400	700		
Phreatic ET		<300	<1,000	2,000	
Groundwater withdrawals		<300			<sup>3</sup> <300
Incidental recharge	<300				
Totals	1,600	1,600			
Storage change	Between -300 and 0				
Total storage (maf)	4.5 maf				
					<sup>4</sup> 1.48–3.94 maf

<sup>1</sup>Predevelopment. Values obtained from unpublished tabular data. Plates in report show only qualitative ranges of values.

<sup>2</sup>Based on re-running Freethy and Anderson (1986) models with revised precipitation data.

<sup>3</sup>Tadayon (2005), value as of year 2000.

<sup>4</sup>Mason and others (2007).



**Figure 9.** Water budget for Detrital Valley showing recent (2007–08) human-driven effects and long-term (1940–2008) natural effects, Mohave County, Arizona. All units are acre-feet per year, unless otherwise indicated; maf, millions of acre-feet.

walls (Laney, 1979), or as focused discharge from high-permeability sediments. Bank-storage effects probably extend no more than half a mile from the lake's edge (Bales and Laney, 1992; McBride and Pfannkuch, 1975) and have been observed in wells adjacent to Lake Mead where water levels rise and fall in response to Lake Mead water-surface elevation changes (Laney, 1979).

Monkey Cove Spring, located at 1,150 ft above sea level, is the largest identified spring below Lake Mead's maximum water-surface elevation; its discharge has been measured to be as much as 1,700 gallons per minute (gal/min). From 1971 to April 2003, this spring was submerged by as much as 75 feet of water from Lake Mead. From April 2003 through 2009, the spring was above the reservoir water level and discharged water visibly. The spring's discharge was 1,350 gal/min in February 2008, but only 380 gal/min in May 2009. Many factors can contribute to such large changes in measured discharge, and future studies could investigate this further. Because the source of water for Monkey Cove and other similar springs was not clear, flow rates of these springs were not used to estimate natural discharge.

## Groundwater Withdrawals

Groundwater withdrawals from the Detrital Valley alluvial aquifer (2007–08) were less than 300 acre-ft/yr. This total is reasonable considering that about 1,700 people lived in the valley in 2007 and there were no major industrial or agricultural water uses. Water use was evenly divided between the SSD population (700 people) and the population served by CWSs (1,000 people). CWS water use occurred in and around the community of Dolan Springs.

## Incidental Recharge

Incidental recharge to the Detrital Valley alluvial aquifer (2007–08) was less than 300 acre-ft/yr, which reflects the low rate of water use in the valley. Individual components of incidental recharge considered by this report (water-supply pipes, septic systems, and WWTPs) were each less than 100 acre-ft/yr. There are no WWTPs in Detrital Valley.

## Groundwater Storage

Groundwater storage in the Detrital Valley alluvial aquifer was estimated to be decreasing during 2007–08, but at less than 300 acre-ft/yr. From 1940 through 2008, between 3,400 and 15,000 acre-ft of water have been withdrawn. Total groundwater storage to a depth of 1,300 ft bls is estimated to be 4.5 maf (M. Truini, U.S. Geological Survey, written commun., 2011).

# Hualapai Valley Water Budget

## Natural Recharge

Natural recharge to the Hualapai Valley alluvial aquifer is estimated to be 5,700 acre-ft/yr (fig. 10; table 2), on the basis of BCM calculations. Hualapai Valley receives an average of 610,000 acre-ft/yr (9.4 in/yr) of precipitation. Of this total precipitation, 384,000 acre-ft/yr falls on the valley floor, and 226,000 acre-ft/yr falls on mountain blocks. All valley-floor precipitation returns to the atmosphere through ET, and about 4,800 acre-ft/yr (2.1 percent) of mountain-block precipitation eventually becomes recharge. Total natural recharge comprises 4,400 acre-ft/yr of in-place recharge from mountain blocks, 400 acre-ft/yr of recharge resulting from mountain-block runoff, 600 acre-ft/yr from streamflow in Truxton Wash, and 300 acre-ft/yr as underflow beneath Truxton Wash.

Estimates of underflow into Hualapai Valley beneath Truxton Wash are complicated by large changes in water-table elevation since the 1950s. Under predevelopment conditions, an estimated 300 acre-ft/yr entered Hualapai Valley as underflow on the basis of water-table maps (Freethy and Anderson, 1986). During the 1960s, however, groundwater gradients appeared reverse, resulting in about 1,000 acre-ft/yr flowing out of Hualapai Valley (Gillespie and Bentley, 1971). By the early 1980s, groundwater levels had risen and underflow might have once again flowed into Hualapai Valley, although limited data make this unclear. Well B-23-13 19DCB exemplifies this pattern: in 1944, the depth-to-water was 53 ft; in 1973, it was 111 ft; in 1980, it was 61 ft; and in 2001, it was 98 ft. For the purposes of this report, with the understanding that the value is highly uncertain, underflow into Hualapai Valley was estimated to be 300 acre-ft/yr—the same value reported by Freethy and Anderson (1986).

## Natural Discharge

Natural discharge from the Hualapai Valley alluvial aquifer is estimated to be 5,700 acre-ft/yr. This value was assumed to be the same as estimated natural recharge, reflecting a key assumption discussed further in the “Effects of Groundwater Withdrawals” section. Although most natural discharge occurs as groundwater discharge to Lake Mead, some unquantified amount could occur as phreatic ET.

Tillman and others (2011) estimated that between 1,800 and 3,200 acre-ft/yr of groundwater discharges from the Hualapai Valley alluvial aquifer through phreatic ET, but that value was not used in this water budget. The rationale was the same as that used for Detrital Valley (see the “Natural Discharge” subsection within the “Detrital Valley Water Budget” section).

### Groundwater Withdrawals

Groundwater withdrawals from the Hualapai Valley alluvial aquifer (2007–08) were about 9,800 acre-ft/yr, of which 1,200 acre-ft/yr was delivered to Sacramento Valley through an interbasin transfer. Of the 8,600 acre-ft/yr used in Hualapai Valley, essentially all was for municipal and domestic uses. The City of Kingman municipal utility served about 84 percent of all residents (33,000 people) in Hualapai Valley and, in so doing, used about 7,600 acre-ft/yr of water. CWSs pumped about 500 acre-ft/yr, serving about 7 percent of the valley population (2,800 people). About 9 percent of the valley population (3,500 people) is estimated to be SSD, withdrawing about 500 acre-ft/yr.

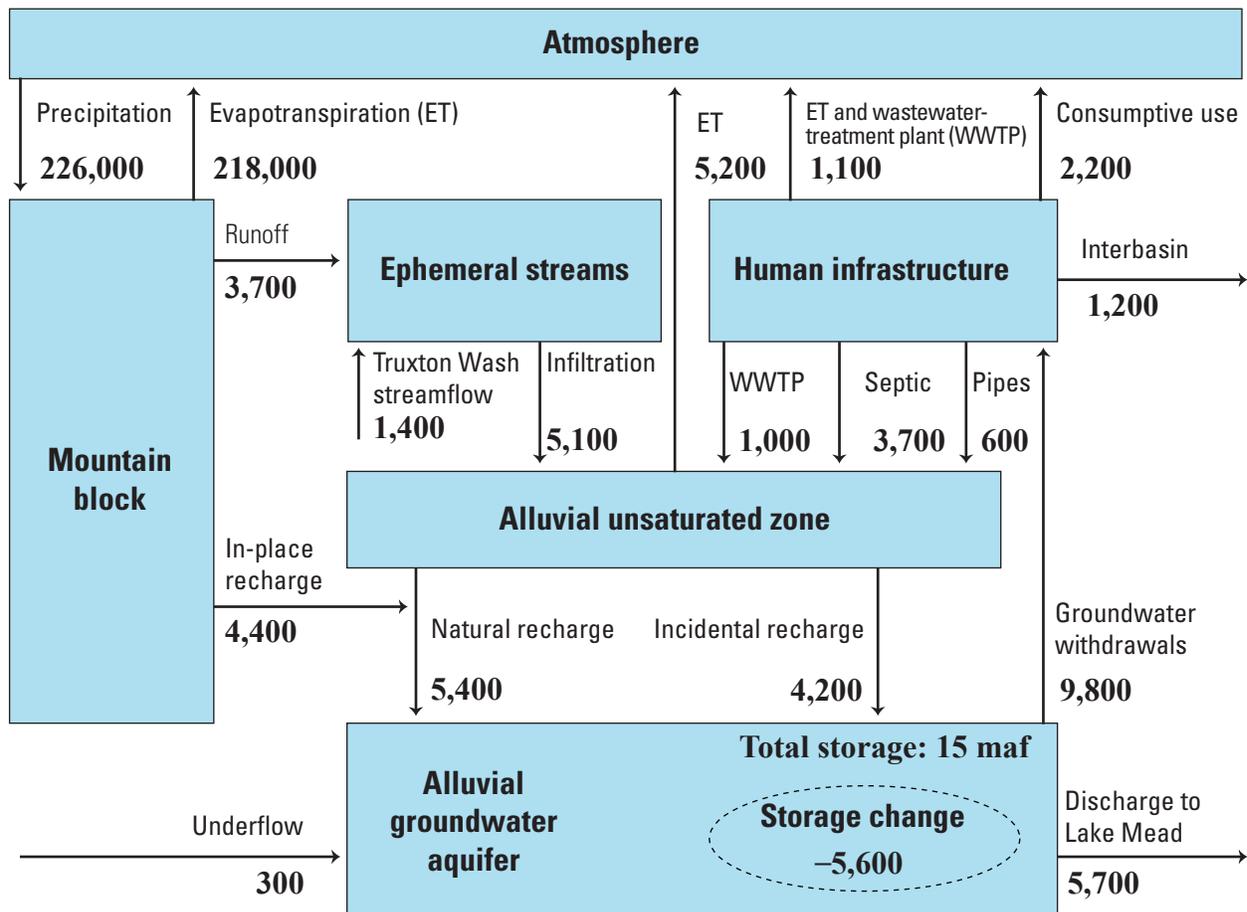
About 2 acre-ft/yr was reported to be withdrawn for industrial use in Hualapai Valley (2007–08). This value is considered negligible in the context of this report’s water budgets.

### Incidental Recharge

About 4,200 acre-ft/yr of water (2007–08) recharged the Hualapai Valley alluvial aquifer as a result of human infrastructure and water use. About 500 acre-ft/yr of this was from leaks in water-supply lines, 3,000 acre-ft/yr from septic-system drain fields, and 800 acre-ft/yr from the rapid infiltration basins of the Hualapai WWTP. Although about 80 percent of septic-system discharge recharged the aquifer, only 38 percent of effluent treated by the Hualapai Valley WWTP did so. This difference is due to high summertime ET from the WWTP wetlands.

The Hualapai WWTP reported treating about 2,000 acre-ft/yr of effluent, but 5,100 acre-ft/yr of effluent is estimated to be produced by residents of City of Kingman. The discrepancy is caused by about 60 percent of City of Kingman residents in Hualapai Valley receiving municipal water, but discharging their effluent to septic systems. This 60-percent value is similar to the 50-percent value estimated by the City of Kingman municipal utility (J. Corwin, City of Kingman, oral commun., 2010).

## Hualapai Valley Water Budget



**Figure 10.** Water budget for Hualapai Valley showing recent (2007–08) human-driven effects and long-term (1940–2008) natural effects, Mohave County, Arizona. All units are acre-feet per year, unless otherwise indicated; maf, millions of acre-feet.

**Table 2. Groundwater-budget values for Hualapai Valley, Mohave County, Arizona; includes values from previous studies for comparison.**

[acre-ft/yr, acre-feet per year; maf, millions of acre-feet; <, less than; WWTP, wastewater-treatment plant; CWS, community water supplier; SSD, self-supplied domestic; ET, evapotranspiration]

Water-budget component	This study		Other studies			
	Inflow to aquifer (acre-ft/yr)	Outflow from aquifer (acre-ft/yr)	Gillespie and Bentley (1971) (acre-ft/yr)	Freethey and Anderson (1986) <sup>1</sup> (acre-ft/yr)	Tillman and others (2011) (acre-ft/yr)	Other (see footnote <sup>5</sup> ) (acre-ft/yr)
Natural recharge	5,700		5,000	3,500	24,000	<sup>3</sup> 2,000–2,500
Mountain-block recharge	4,400					
Truxton Wash stream-channel recharge	600					
Other ephemeral stream-channel recharge	400					
Underflow at Truxton Wash	<sup>4</sup> 300		<sup>5</sup> -1,000	300		
Natural discharge		5,700	5,000			
to Lake Mead		5,700	5,000	3,800		
Phreatic ET		<300		<1,000	3,000	
Groundwater withdrawals		9,800				<sup>6</sup> 8,200
City of Kingman municipal		7,600				
CWS		500				
SSD		500				
Interbasin transfer		1,200				
Incidental recharge	4,200					
Infrastructure leakage	500					
Septic systems	3,000					
Treated WWTP effluent	800					
Totals	9,900	15,500				
Storage change		-5,600				
Total storage (maf)		5 maf	10.5–21 maf			<sup>3</sup> 3.8–10.1 maf

<sup>1</sup>Predevelopment. Values obtained from unpublished tabular data. Plates in report show only qualitative ranges of values.

<sup>2</sup>Based on re-running Freethey and Anderson (1986) models with revised precipitation data.

<sup>3</sup>Ivanich and Conway (2007).

<sup>4</sup>From Freethey and Anderson (1986) pre-development conditions, as there are insufficient data to calculate a current-condition value.

<sup>5</sup>Negative indicates basin outflow, rather than inflow, because of large Truxton Canyon pumping withdrawals in the 1960s.

<sup>6</sup>Tadayon (2005), value as of year 2000.

## Groundwater Storage

Groundwater storage was estimated to be decreasing at about 5,600 acre-ft/yr during 2007–08. While 9,800 acre-ft/yr was withdrawn by pumping, about 4,200 acre-ft/yr returned to the aquifer as incidental recharge—about 43 percent of withdrawals. One factor to consider in evaluating this number, however, is that 1,200 acre-ft/yr of water left Hualapai Valley through an interbasin transfer; excluding this interbasin-transfer water, the ratio of incidental recharge to withdrawals is about 50 percent. Total groundwater storage to a depth of 1,300 ft bls is estimated to be 15 maf, reflecting 2006 water-level conditions (M. Truini, U.S. Geological Survey, written commun., 2011).

## Sacramento Valley Water Budget

### Natural Recharge

Natural recharge to the Sacramento Valley alluvial aquifer is estimated to be 6,000 acre-ft/yr (fig. 11; table 3), on the basis of BCM calculations. Sacramento Valley receives an average of 770,000 acre-ft/yr (9.1 in/yr) of precipitation. Of this total precipitation, 440,000 acre-ft/yr falls on the valley floor, and 340,000 acre-ft/yr falls on mountain blocks. All valley-floor precipitation returns to the atmosphere through ET, and about 6,000 acre-ft/yr (1.8 percent) of mountain-block precipitation becomes recharge. This recharge comprises

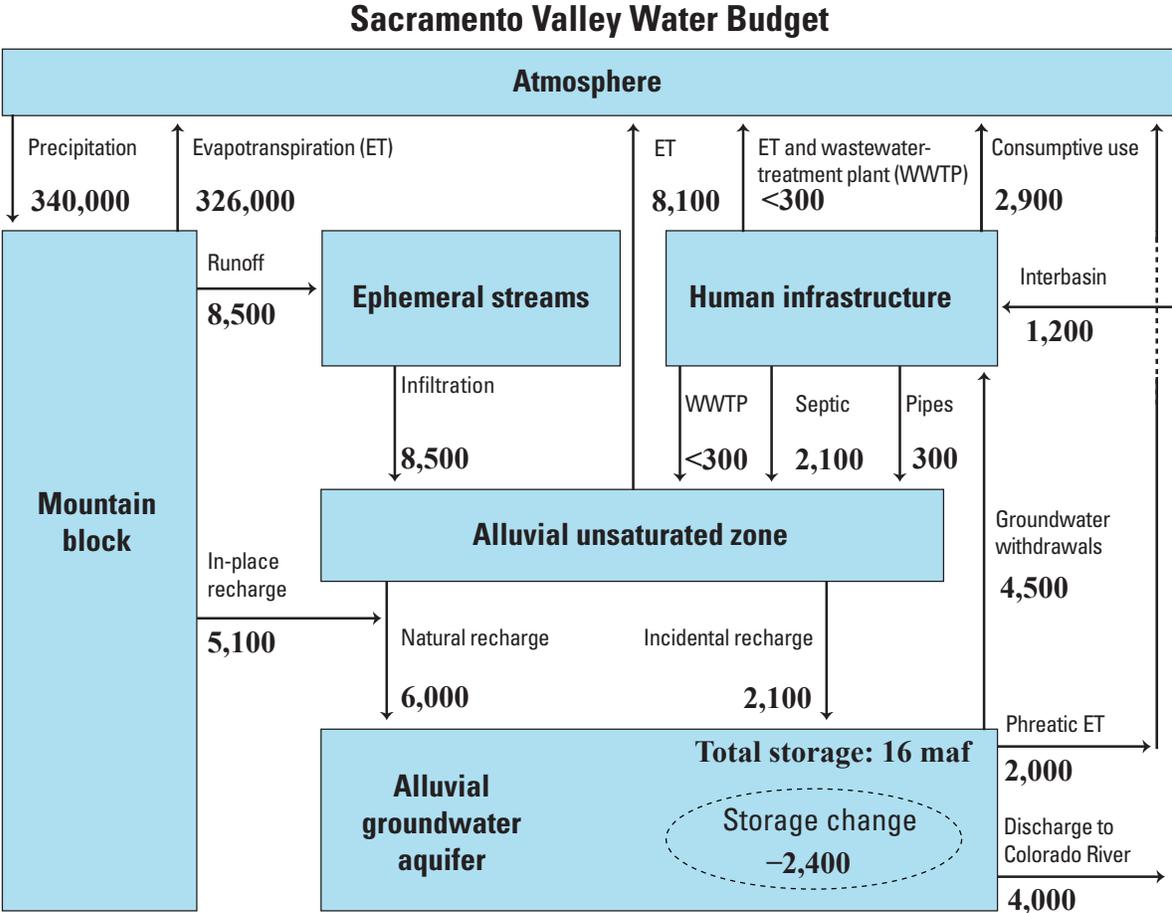
5,100 acre-ft/yr of in-place recharge and 800 acre-ft/yr of recharge resulting from runoff.

The stream-channel network in southernmost Sacramento Valley drains southward to Bill Williams Valley, even though groundwater in the area flows northward and remains in Sacramento Valley (Anning and others, 2007; fig. 1). Some surface-water runoff, therefore, exits the study area through Mohave and Castaneda Washes—perhaps as much as 700 acre-ft/yr, as indicated by the BCM. However, even if all of this runoff exited the study area (unlikely, as some probably would infiltrate), under the assumption that 10 percent of BCM runoff becomes recharge, total recharge would be reduced by only 70 acre-ft/yr (about 1 percent). Because of uncertainty in these quantities, for the purposes of this report, streamflow to Bill Williams Valley is not included in water-budget calculations. Streamflow monitoring on Mohave and Castaneda Washes would aid future studies in better understanding the hydrology of this area.

### Natural Discharge

Natural discharge from the Sacramento Valley alluvial aquifer is estimated to be 6,000 acre-ft/yr. This value is assumed to be the same as estimated natural recharge, reflecting a key assumption discussed further in the “Effects of Groundwater Withdrawals” section. Natural discharge occurs mostly as groundwater discharge to the Colorado River, but some could occur through phreatic ET near Havasu NWR. As discussed in the preceding section, amounts of ephemeral streamflow to Bill Williams Valley are uncertain and were not included in this total natural discharge value.

Tillman and others (2011) estimated that between 5,800 and 10,000 acre-ft/yr of groundwater discharges from the Sacramento Valley alluvial aquifer through phreatic ET, but spatial analysis and field verification of their data suggest a more reasonable value would be 2,000 acre-ft/yr. Their analysis indicated between



**Figure 11.** Water budget for Sacramento Valley showing recent (2007–08) human-driven effects and long-term (1940–2008) natural effects, Mohave County, Arizona. All units are acre-feet per year, unless otherwise indicated; maf, millions of acre-feet.

**Table 3. Groundwater-budget values for Sacramento Valley, Mohave County, Arizona. Includes values from previous studies for comparison.**

[acre-ft/yr, acre-feet per year; maf, millions of acre-feet; ≤, less than or equal to; WWTP, wastewater-treatment plant; CWS, community water supplier; SSD, self-supplied domestic; ET, evapotranspiration]

Water-budget component	This study		Other studies			
	Inflow to aquifer (acre-ft/yr)	Outflow from aquifer (acre-ft/yr)	Gillespie and Bentley (1971) (acre-ft/yr)	Freethey and Anderson (1986) <sup>1</sup> (acre-ft/yr)	Tillman and others (2011) (acre-ft/yr)	Other (see footnote <sup>5</sup> ) (acre-ft/yr)
Natural recharge	6,000		4,000	4,000	<sup>2</sup> 6,000	<sup>3</sup> 1,000; <sup>4</sup> 12,300
Mountain-block recharge	5,200					
Ephemeral stream-channel recharge	800					
Natural discharge to Colorado River		6,000	4,000	4,000		<sup>4</sup> 12,300
Phreatic ET		<sup>5</sup> 4,000	4,000	4,000		
Groundwater withdrawals		<sup>5</sup> 2,000		<1,000	<sup>6</sup> 10,000	
City of Kingman municipal		4,500				<sup>7</sup> 2,150
CWS		500				
SSD		2,000				
Industrial		100				
Incidental Recharge	<sup>8</sup> 2,100					
Infrastructure leakage	<sup>8</sup> <300					
Septic systems	<sup>8</sup> 1,700					
Treated WWTP effluent	<sup>8</sup> <300					
Totals	8,200	10,500				
Storage change		-2,400				
Total storage (maf)		16 maf	6.5–13 maf			<sup>9</sup> 3.6–9.5 maf

<sup>1</sup>Predevelopment. Values obtained from unpublished tabular data. Plates in report show only qualitative ranges of values.

<sup>2</sup>Based on re-running Freethey and Anderson (1986) models with revised precipitation data.

<sup>3</sup>Rascona (1991).

<sup>4</sup>Metzger and Loeltz (1973).

<sup>5</sup>Partitioning between Colorado River and phreatic ET uncertain because of a lack of data.

<sup>6</sup>Only 2,000 acre-ft/yr of this value occurs in zones with known woody and herbaceous plants.

<sup>7</sup>Tadayon (2005), value as of year 2000.

<sup>8</sup>Includes the effects of 1,200 acre-ft/yr of water transferred from Hualapai Valley for City of Kingman.

<sup>9</sup>Conway and Ivanich (2007).

3,800 and 7,000 acre-ft/yr of phreatic ET occurring within and along ephemeral stream channels in the valley; however, field verification of vegetation in these channels found no evidence of phreatophytes except possibly mesquite, which also can survive in nonphreatic conditions. Depths-to-water in wells near these stream channels range from 62 to more than 800 ft (Anning and others, 2007). Because of the apparent “greenness” of these stream channels, the regional regression model used by Tillman and others appears to overestimate phreatic ET in these stream channels.

Tillman and others (2011) also indicated that about 2,000 acre-ft/yr of phreatic ET occurs in a zone of known woody/herbaceous plants at the downstream end of Sacramento Wash. This zone coincides with a 700-acre stand of saltcedar (fig. 12) within Havasu NWR (see fig. 1). Saltcedar is an invasive plant that can obtain water from the groundwater system (Di Tomaso, 1998). Dividing Tillman’s estimated 2,000 acre-ft/yr by 700 acres produces 3.1 ft/yr of water use, which is within a published range (2.5–4 ft/yr) for saltcedar-dominated riparian vegetation areas (Dahm and others, 2002).

Groundwater hydrology within Havasu NWR is complicated by diversions and dams associated with Topock Marsh (Guay and others, 2006; Guay, 2001). Groundwater used by saltcedar could be obtained from the Sacramento Valley alluvial aquifer, groundwater recharge from the Colorado River, or both. Additional data, particularly groundwater levels within and around Havasu NWR, would help to better understand the hydrology of this area. For the purposes of this water budget—and with the understanding that this value is uncertain—2,000 acre-ft/yr is estimated to return to the atmosphere from the Sacramento Valley alluvial aquifer through phreatic ET.

## Groundwater Withdrawals

Groundwater withdrawals from the Sacramento Valley alluvial aquifer (2007–08) were about 4,500 acre-ft/yr, and an additional 1,200 acre-ft/yr was transferred into the valley from Hualapai Valley by the City of Kingman. About

two-thirds of these withdrawals were used for municipal and domestic purposes; the remaining was used for industrial processes, primarily thermoelectric power generation. The City of Kingman required about 1,700 acre-ft/yr of water to meet customer demand—500 acre-ft/yr by withdrawal from Sacramento Valley and 1,200 acre-ft/yr by a transfer from Hualapai Valley. CWSs withdrew about 2,000 acre-ft/yr, serving about 55 percent of the Sacramento Valley population. About 4 percent of the population (800 people) is estimated to be SSD, withdrawing less than 300 acre-ft/yr.

## Incidental Recharge

About 2,100 acre-ft/yr of water (2007–08) recharged the Sacramento Valley alluvial aquifer as a result of human infrastructure and water use. Less than 300 acre-ft/yr of this recharge was from leaks in water-supply lines, about 1,700 acre-ft/yr from septic-system drain fields, and less than 300 acre-ft/yr from treated effluent discharging from the City of Kingman WWTP in Sacramento Valley.

The WWTP in Sacramento Valley treated about 400 acre-ft/yr of effluent, but 1,100 acre-ft/yr of effluent was estimated to be produced by City of Kingman residents. As in Hualapai Valley, this discrepancy was explained by 60 percent of City of Kingman residents receiving municipal water but discharging effluent to septic systems instead of the WWTP. Three additional WWTPs in Sacramento Valley (Arizona Department of Water Resources, 2009b, p. 43) likely treated less than a combined 100 acre-ft/yr of effluent, an amount considered to be negligible.

Almost two-thirds of water withdrawn from the Sacramento Valley alluvial aquifer (2007–08) was used consumptively, which is a larger proportion than is used by the other valleys in the study area. The reason for this was that there was a much larger amount of industrial water use in Sacramento Valley. Thermoelectric-power generation was one of the dominant industrial processes, and consumptive use occurs through two general pathways (World Nuclear Association, 2010). First, about 1 acre-ft of water evaporates from cooling towers for every 500 megawatt-hours of electricity produced. Second, blowdown water is discharged to offset the salinity buildup caused by cooling-tower evaporation; blowdown water enters lined evaporation ponds and returns to the atmosphere through evaporation.

## Groundwater Storage

Groundwater storage was estimated to be decreasing at about 2,400 acre-ft/yr in the Sacramento Valley alluvial aquifer during 2007–08. While 4,500 acre-ft/yr was being withdrawn by pumping, about 2,100 acre-ft/yr was returning as incidental recharge. This suggests that about half of all withdrawn water returned as incidental recharge. However, about 30 percent of this incidental recharge was associated with an interbasin transfer from Hualapai Valley; excluding interbasin transfers, the

ratio of incidental recharge to withdrawals is about 34 percent. Total groundwater storage to a depth of 1,300 ft bls is estimated to be 16 maf, reflecting 2006 water-level conditions (M. Truini, U.S. Geological Survey, written commun., 2011).

## Discussion

### Comparison with Other Studies

Water budgets in this report generally contained higher natural recharge and discharge values than previously published values. In light of the varied approaches taken and the inherent uncertainty, estimated values in this report still are considered to be comparable to those of most previous reports. A summary of values from tables 1, 2, and 3 is given here.

Freethy and Anderson's (1986) values for natural recharge and discharge are 33–50 percent less than those of this report. Their method for recharge-estimation used a linear equation relating precipitation above 8 in/yr to recharge across all alluvial basins of Arizona. Precipitation values used in the Freethy and Anderson (1986) recharge equation were estimated by using generalized, hand-contoured maps.

Natural-recharge values from Tillman and others (2011), estimated using the BCM, agree closely with values from this report. Tillman and others also re-applied the model of Freethy and Anderson (1986) using PRISM estimates of precipitation instead of hand-contoured precipitation maps; the resulting recharge estimates are higher than values reported by Freethy and Anderson (1986) and are similar to the values of this report. As discussed elsewhere, estimates of phreatic ET by Tillman and others (2011) generally are considered to be



**Figure 12.** Stand of riparian vegetation dominated by saltcedar, looking upstream in Sacramento Wash, Havasu National Wildlife Refuge, Mohave County, Arizona.

high for the study area in light of the additional spatial analyses and field verification described in this report.

Gillespie and Bentley's (1971) estimates of natural discharge from Hualapai and Sacramento Valleys are 24 and 33 percent less, respectively, than those of this report. Another difference between Gillespie and Bentley's report and this one was underflow between Hualapai Valley and Truxton Canyon. They estimated 1,000 acre-ft/yr of underflow exiting Hualapai Valley in the 1970s, whereas this report estimates 300 acre-ft/yr of underflow entering Hualapai Valley in 2007–08. This discrepancy is discussed in the "Natural Recharge" subsection of the "Hualapai Valley Water Budget" section.

Two studies of Sacramento Valley reported values that disagree with those of this report. Rascona (1991) estimated natural groundwater discharge at 1,000 acre-ft/yr, four to six times less than other published values. However, Rascona subtracted groundwater withdrawals from natural discharge, an approach not taken in this report because groundwater withdrawals were assumed to have not yet depleted natural discharge or streamflow. Metzger and Loeltz (1973) estimated natural recharge and discharge each at 12,300 acre-ft/yr by applying an empirical relation between precipitation and recharge developed for central Nevada (Eakin and others, 1951, p. 79–81). This value is more than twice as large as any other published recharge or discharge value for Sacramento Valley.

## Effects of Groundwater Withdrawals

From 1940 through 2008, total groundwater withdrawals are about 3,400 acre-ft for Detrital Valley, 270,000 acre-ft for Hualapai Valley, and 170,000 acre-ft for Sacramento Valley (fig. 13; appendix 1). These values are general estimates, as they do not include estimates of unreported withdrawals prior to 2007 and are based on historical data that contain uncertainty. These total withdrawals are low compared to those of other Arizona groundwater basins (for example, 134,000 acre-ft/yr is withdrawn from the Willcox Basin; Tadayon, 2005), but they are significant for the study area given the low natural-recharge rates.

Any amount of groundwater withdrawn by pumping is accounted for by some combination of decrease in groundwater storage, reduction in natural discharge, and/or increase in natural recharge (Theis, 1940; Leake and others, 2008). Determining the distribution among these three effects typically is difficult without a groundwater-flow model. However, because of the locations of withdrawals to date in the study area and the large storage volumes of the alluvial aquifers, a reasonable assumption could be that all, or nearly all water withdrawn to date has been accounted for by storage depletion. The main supporting evidence for this assumption is that water-level declines, as of 2006, have not yet reached downgradient discharge areas (Anning and others, 2007). Mountain-front recharge in the study area is unlikely to be affected by withdrawals, because water tables are disconnected from mountain-front stream channels, and their lowering would not be able to induce additional recharge (Anderson, 1995).

Leake and others (2008) estimated that, after 100 years of continuous groundwater withdrawals near Dolan Springs in Detrital Valley, between 1 and 50 percent of withdrawn water would be accounted for by a reduction in natural discharge to and/or increase in recharge from Lake Mead. This estimate is based on a simplified groundwater-flow model with a range of transmissivity values. Ultimately, there are only two possible long-term outcomes: (1) long-term groundwater withdrawals eventually (perhaps after centuries) obtain 100 percent of their water from reduction in discharge and/or depletion of streamflow with no additional long-term storage depletion (Leake and others, 2008), or (2) long-term groundwater withdrawals are greater than all discharge or capturable streamflow, and the aquifer becomes depleted of all recoverable water. Estimating the time required for either of these ultimate outcomes to occur is beyond the scope of this report.

Continued withdrawals from the Kingman subbasin of Hualapai Valley will continue to lower water levels and steepen gradients, but such changes may be more pronounced because of the geometry of this subbasin. The Kingman subbasin is connected hydrologically to the adjacent Hualapai subbasin by alluvial fill that is only 300–600 ft deep (fig. 3), suggesting limited connectivity between the Kingman subbasin and other areas of Hualapai Valley. As water is withdrawn from the Kingman subbasin, water from other Hualapai Valley subbasins would flow preferentially toward the Kingman subbasin, but would be limited by the low transmissivity along its northern boundary.

## Incidental Recharge and Water Quality

Across the study area, incidental recharge accounted for about 35 percent of total recharge, which could raise water-quality concerns (Asano and Cotruvo, 2004). Although incidental-recharge rates are small relative to total aquifer storage, localized water-quality issues might arise if wells are near areas of incidental recharge. Arizona regulatory agencies already have expressed concern about high groundwater nitrate concentrations in the Kingman subbasin of Hualapai Valley. Natural nitrate accumulations in the unsaturated zone can be mobilized by incidental recharge and result in high groundwater nitrate concentrations (Walvoord and others, 2003), however, no studies have been done to investigate this issue in the study area (J. Corwin, City of Kingman, oral commun., 2010). Incidental recharge from leaking water-supply pipes also can cause water-quality concerns. Disinfection by-products, such as trichloromethane, can be produced as drinking-water disinfectants interact with natural organic matter in the aquifer (Izbicki and others, 2010; Mahler and others, 2006). No studies of this process have been undertaken in the study area.

### Limitations and Uncertainty in Water Budgets

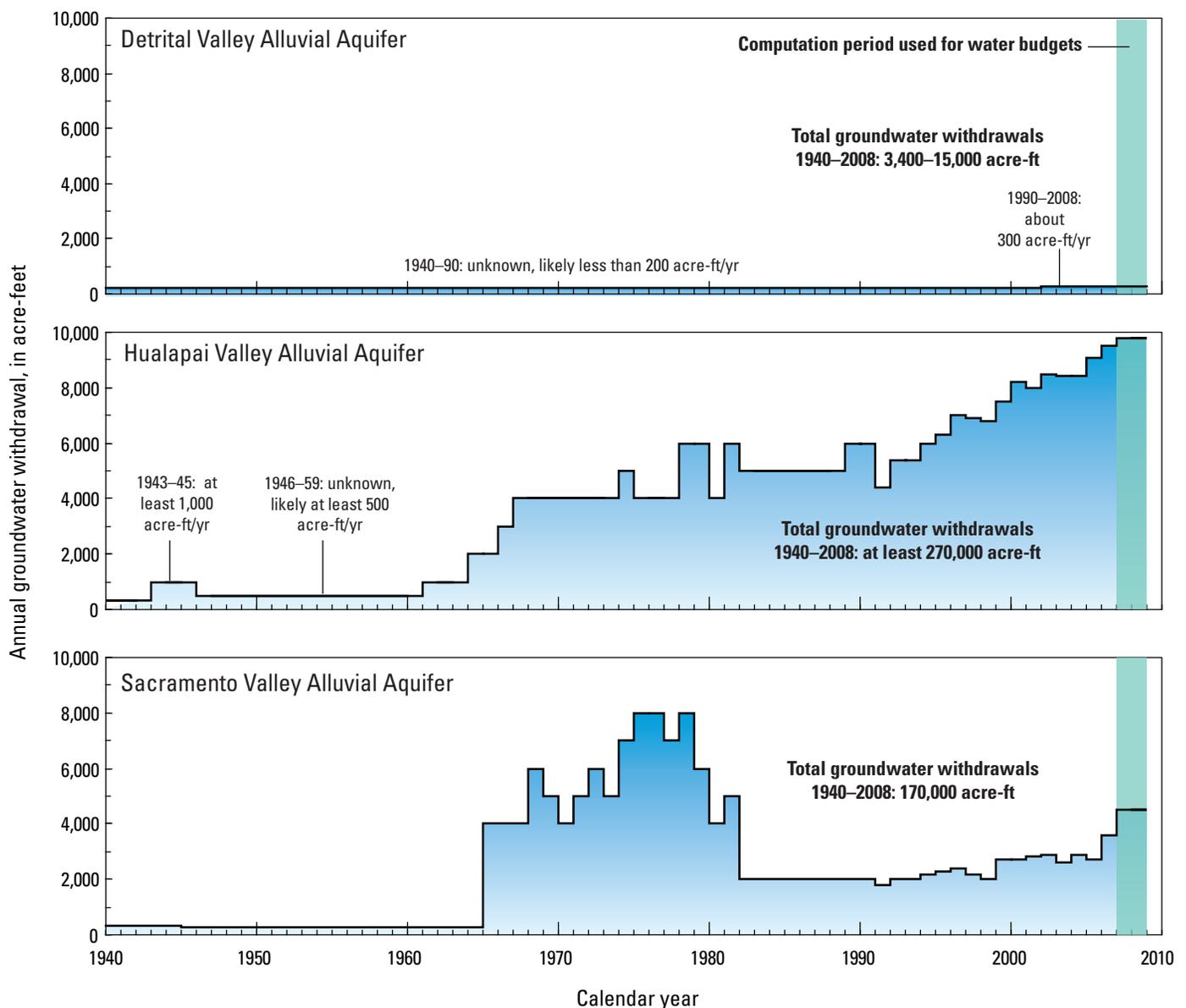
Water budgets are “lumped-parameter models,” in that their values present average conditions both in space and over time. The reality is far more complex—groundwater systems have “complex, time-varying, three-dimensional flow systems” (Upper San Pedro Partnership, 2007, p. 41). This can limit the applicability of water budgets for answering detailed questions about groundwater systems, and groundwater-flow models could be required in those cases.

Uncertainty and sources of error often are not discussed at great length in water budgets, leading to a false sense of

confidence in their values (Winter, 1981). Although a rigorous, quantitative estimate of uncertainty is beyond the scope of this report, qualitative discussions of limitations and major sources of uncertainty for natural recharge, natural discharge, incidental recharge, and storage are provided in this section.

### Natural Recharge

Natural recharge values in this report are subject to uncertainty inherent in the BCM. Precipitation data used by the BCM (PRISM) can have absolute errors of more than 20 percent on an annual basis, particularly in mountainous areas of



**Figure 13.** Groundwater withdrawals from Detrital, Hualapai, and Sacramento Valley alluvial aquifers, 1940-2008, Mohave County, Arizona.

the western United States (Daly and others, 2008). BCM data for hydraulic properties of bedrock and soil are extrapolated to large regional areas on the basis of generalized geologic maps and have not been field-verified within this study area. Potential ET is estimated only as a single set of 12 monthly values in the BCM, instead of as a time-series dataset, as for precipitation.

Another consideration is that the BCM runs at a monthly time-step and, consequently, cannot model intense, short-duration thunderstorms. Such thunderstorms can produce localized precipitation that, on the scale of hours, exceeds potential ET and becomes recharge. On a monthly time-scale, however, cumulative precipitation from small thunderstorms is much less than total potential ET. The amount of recharge occurring in this manner and not accurately modeled by the BCM is not known for the study area.

For this report, 10 percent of BCM runoff eventually was assumed to become recharge. This single value encapsulates a number of processes and variability, including location, lithology, antecedent conditions, precipitation intensity, and season. Vegetation also affects this value, as can upward movement of unsaturated-zone water caused by thermal gradients (de Vries and Simmers, 2002). Additional field studies—or possibly a calibrated groundwater-flow model—would be needed to better constrain this value.

The period 1940–2008 was used to estimate natural recharge, but it might not fully capture true average conditions. A 1904 winter flood in Truxton Canyon had a peak discharge of 49,000 ft<sup>3</sup>/s (Murphy, 1905). A September 1939 storm produced a peak discharge of 15,000 ft<sup>3</sup>/s in Sacramento Wash near the Colorado River (Metzger and Loeltz, 1973, p. 18). These and other rare flood events indicate that a longer period of record might produce different recharge estimates.

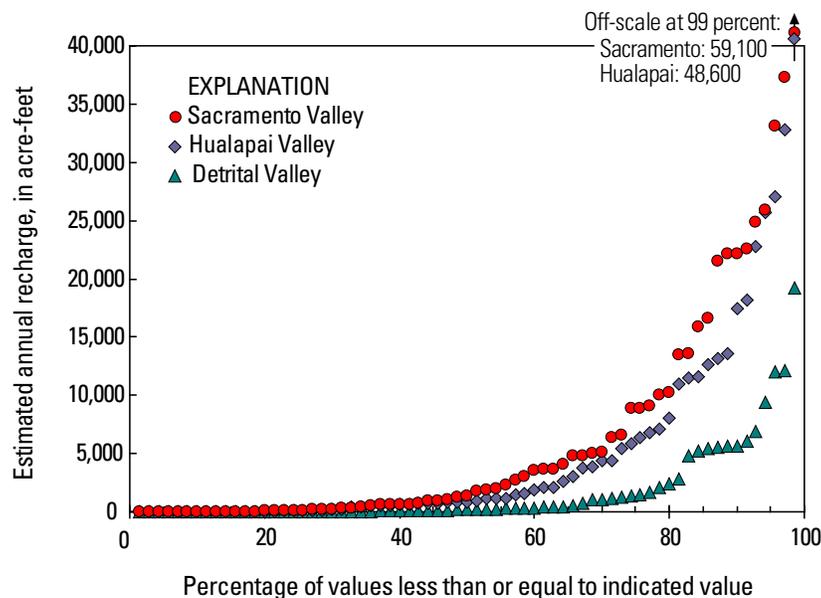
Although water budgets in this report present natural recharge as a single estimate of long-term, mean-annual

recharge, natural variability of recharge is important to remember (de Vries and Simmers, 2002). Natural recharge’s interannual variability spans more than four orders of magnitude within the study area (fig. 14), a characteristic common to the arid southwest. For about 30 percent of the years 1940–2008, annual recharge to Hualapai and Sacramento Valleys was less than 300 acre-ft; for Detrital Valley, 60 percent of years 1940–2008 had annual recharge of less than 300 acre-ft. At the other extreme, the three years that produced the greatest amount of natural recharge (1993, 2005, and 1978) were 5–8 times greater than the mean-annual recharge values shown in this report.

### Natural Discharge

As stated throughout this report, natural discharge is assumed to be equal to natural recharge even though groundwater withdrawals have occurred since about 1940. This assumption is considered reasonable for 2007–08 conditions, but at some point after 2008 it will no longer be valid.

Direct measurements of natural discharge would be ideal, but are not practical for several reasons. Lake Mead has inundated most historical springs (Longwell, 1936) that presumably discharged groundwater from the alluvial aquifers of Detrital and Hualapai Valleys. Impoundments of and diversions from the Colorado River near Havasu NWR have changed the groundwater-discharge regime of the Sacramento Valley alluvial aquifer. Finally, groundwater discharge to the Colorado River in the study area is more than four orders of magnitude less than average Colorado River discharge. A method capable of detecting less than a 1 part in 1,000 change in streamflow presently is not practical.



**Figure 14.** Frequency-duration graphs of annual recharge estimated by the Basin Characterization Model for Detrital, Hualapai, and Sacramento Valley alluvial aquifers, 1940–2008, Mohave County, Arizona.

## Groundwater Withdrawals

Groundwater withdrawals by municipal, industrial, and CWSs estimated in this report are subject to uncertainties inherent in these data, including inaccurate flowmeters and incomplete or erroneous reporting of withdrawals. Withdrawals by the SSD population were estimated by using census data; part-year, winter-only residents cause additional uncertainty because they might or might not be counted in the census.

## Incidental Recharge

Incidental-recharge values in this report relied on generalized percentages from studies in similar areas. Many additional processes, including lawn over watering, impervious-cover changes, stream-channel lining, and soil disturbance can affect incidental recharge rates (Scanlon and others, 2006; Shuster and others, 2005) but were not considered. Another factor to consider is the accuracy of groundwater-withdrawal data reported to State agencies because incidental recharge rates are directly proportional to groundwater withdrawal rates. Finally, part-year, winter-only residents likely cause additional uncertainty because they would use more water in the winter than summer and, therefore, could cause consumptive use to be underestimated.

## Total Groundwater Storage

Total groundwater-storage values can be misleading from a resource-management perspective, because they likely do not reflect the actual amount of practically recoverable water (Alley, 2007). In a scenario where groundwater storage is being continually depleted, removal of all groundwater from storage by pumping is impossible from a practical standpoint. Land subsidence, changes in water quality, and increased pumping costs likely would be encountered long before the full reservoir of groundwater storage was depleted.

## Summary

The USGS, in cooperation with ADWR, initiated a study of the hydrogeology and water resources of Detrital, Hualapai, and Sacramento Valleys in 2005. These valleys primarily have a rural but growing population, and the people and industries in them rely almost entirely on groundwater. Effects of human groundwater use are evident, as indicated by water-level declines of more than 130 ft in some areas. The purpose of this report is to present water budgets for each of the valleys, in order to address fundamental hydrologic questions about water input, output, and storage.

The study area is within the Basin and Range physiographic province. Here, alluvial basins are bounded and underlain by low-permeability igneous and metamorphic rocks (bedrock) and filled with thousands of feet of alluvial fill eroded from surrounding mountain blocks beginning 13–16

million years ago. The climate is arid to semiarid, having average annual precipitation of about 9.2 inches, and no major perennial-stream reaches. Detrital Valley has an area of 900 mi<sup>2</sup>, has groundwater flow from south to north toward Lake Mead, and is sparsely populated. Hualapai Valley has an area of 1,200 mi<sup>2</sup>, has a closed-basin surface-water system draining to Red Lake Playa, but has groundwater flow from south to north toward Lake Mead. Sacramento Valley has an area of 1,600 mi<sup>2</sup>, and groundwater flows generally north to south then west toward the Colorado River.

Water budgets were presented by using a reservoir-and-conveyance approach, wherein water moves between reservoirs (such as the unsaturated zone and the alluvial aquifer) by way of conveyances as a result of natural and human processes (such as infiltration, groundwater withdrawals, and groundwater discharge). Average values for natural processes were estimated using the period from 1940 to 2008, and average values for human processes were estimated from 2007–08 data.

Natural recharge from precipitation was calculated by using a physically based distributed model (the BCM) that calculated recharge and runoff at monthly intervals on the basis of precipitation, climate, soil, and geology hydraulic-property data. Estimated discharge from upland springs was subtracted from recharge and added to runoff, and 10 percent of runoff was assumed to become recharge. Water also entered Hualapai Valley from an adjacent upgradient valley.

Natural discharge is assumed to be equal to natural recharge, reflecting a key assumption that groundwater withdrawals to date have not yet reduced discharge or depleted streamflow. Discharge occurs primarily as groundwater discharge to Lake Mead or the Colorado River. Some discharge also can occur when groundwater is used by vegetation, particularly in Sacramento Valley near Havasu NWR.

Groundwater withdrawals were calculated by compiling data from the City of Kingman, CWSs, industrial facilities, and by estimating withdrawals of the SSD population. Withdrawals by the SSD population were estimated by using census data, because these withdrawals are neither monitored nor reported to State agencies.

Incidental recharge—recharge that occurs as result of human infrastructure and water use—was estimated for leaking water-supply pipes, septic systems, and WWTPs. Twenty-eight percent of groundwater withdrawn for domestic use is estimated not to become recharge because of consumptive use. One-hundred percent of groundwater used for industrial processes is assumed to be consumptively used. Fifty-two percent of water treated at WWTPs is estimated to return to the atmosphere through ET. Of the remaining water, 80 percent is assumed to become recharge.

Annual changes in groundwater storage were calculated as the residuals of inputs subtracted from outputs. Because of assumptions, storage change ultimately reflected only the imbalance between groundwater withdrawals and incidental recharge.

Natural recharge and discharge for the Detrital Valley alluvial aquifer both are about 1,400 acre-ft/yr each. Natural

discharge occurs as groundwater discharge to Lake Mead. Springs below the maximum lake level of Lake Mead might or might not represent discharge from the alluvial aquifer; bank storage from Lake Mead could be the source of some or all of their discharge. Groundwater withdrawals and incidental recharge each were less than 300 acre-ft/yr in 2007–08. Groundwater storage is decreasing at rates less than 300 acre-ft/yr. Total groundwater storage is estimated to be 4.5 maf.

Natural recharge and discharge for the Hualapai Valley alluvial aquifer each are about 5,700 acre-ft/yr. Natural discharge occurs as groundwater discharge to Lake Mead. Of this total recharge, 900 acre-ft/yr is estimated to be from water entering the valley in or beneath Truxton Wash from an upgradient valley; this value is uncertain because of alteration to the natural system in that valley by pumping. Groundwater withdrawals are about 9,800 acre-ft/yr, with 1,200 acre-ft/yr of this being transferred to Sacramento Valley for use in that valley. Most of the remaining 8,600 acre-ft/yr is for domestic and municipal use in the City of Kingman. Incidental recharge was about 4,200 acre-ft/yr in 2007–08. Groundwater storage was decreasing by about 5,600 acre-ft/yr as of 2007–08. Total groundwater storage is estimated to be 15 maf.

Natural recharge and discharge for the Sacramento Valley alluvial aquifer each are about 6,000 acre-ft/yr. Natural discharge mostly occurs as groundwater discharge to the Colorado River, although up to 2,000 acre-ft/yr could occur as phreatic ET near Havasu NWR. Groundwater withdrawals are about 4,500 acre-ft/yr, with an additional 1,200 acre-ft/yr transferred into the valley from Hualapai Valley. Incidental recharge is about 2,100 acre-ft/yr. Consumptive use is higher in Sacramento Valley than in the other valleys in the study area because of a higher amount of industrial water use. Groundwater storage was decreasing at about 2,400 acre-ft/yr as of 2007–08. Total groundwater storage is estimated to be 16 maf.

Water budgets in this report had natural recharge and discharge values that are 24–50 percent greater than most previously published values; this difference is not considered unreasonable given the varied approaches and uncertainty inherent in these calculations. Two previous studies of Sacramento Valley disagreed with values in this report, but the methods used in those reports were based on markedly different assumptions.

From 1940 through 2008, about 3,400 acre-ft, 270,000 acre-ft, and 170,000 acre-ft of water were withdrawn from Detrital, Hualapai, and Sacramento Valley alluvial aquifers, respectively. Through 2008 these withdrawals likely have been derived almost entirely from storage depletion, but at some point after 2008, this will no longer be the case. Withdrawals in the Kingman subbasin of Hualapai Valley could produce more pronounced water-level declines than withdrawals in other areas because of limited transmissivity between this subbasin and adjacent ones.

Incidental recharge is an important recharge component, accounting for about 35 percent of total recharge in the study area as of 2007–08; however, there could be water-quality concerns, such as increased concentrations of nitrate and disinfection by-products as a result of this recharge.

There are many sources of uncertainty for water budgets in this report that should be kept in mind. Natural recharge estimates incorporate uncertainty that is inherent in the BCM, as well as additional uncertainty caused by the BCM monthly time-step and the assumption that 10 percent of runoff becomes recharge. Although natural recharge is presented as a single mean-annual number in this report, it varies considerably from year to year. Incidental recharge estimates relied on generalized percentages from other studies and ignored some possible effects. Total groundwater-storage values in this report should be interpreted carefully because it is unlikely that it is practical to recover all of this groundwater in storage.

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# Appendixes

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# Appendix 1. Historic Groundwater-Withdrawal Data

A summary of available groundwater-withdrawal data for Detrital, Hualapai, and Sacramento Valleys is presented here (table 1.1). Data were compiled from sources that vary in accuracy and reporting limits.

Detrital Valley alluvial aquifer groundwater withdrawals appear to have been always less than 1,000 acre-ft/yr, in support mostly of rural and domestic water uses. Quantitative data

are available beginning in 1991 (Tadayon, 2005), and from 1991 to 2008, withdrawals increased. Withdrawals before 1991 are assumed to be less than the value reported for 1991. Total cumulative groundwater withdrawals from the Detrital Valley alluvial aquifer from 1940 to 2008 are at least 3,400 acre-ft and likely less than 15,000 acre-ft.

Hualapai Valley alluvial aquifer groundwater withdrawals probably first exceeded 1,000 acre-ft/yr from 1943 to 1945 when wells were operated to support an army gunnery school at the City of Kingman Airport (Gillespie and Bentley, 1971). Withdrawal amounts from 1946 to 1960 are unknown but likely are at least 500 acre-ft/yr because one of these wells yielded more than 250 gal/min. By 1961, groundwater withdrawals were consistently

**Table 1.1. Annual groundwater withdrawals for Detrital, Hualapai, and Sacramento Valleys, 1940–2008, Mohave County, Arizona.**

[acre-ft/yr, acre feet per year; <, less than indicated amount; >, greater than indicated amount; (parenthesized value) indicates alternate value obtained from unpublished USGS data]

Year	Detrital Valley (acre-ft/yr)	Hualapai Valley (acre-ft/yr)	Sacramento Valley (acre-ft/yr)	Year	Detrital Valley (acre-ft/yr)	Hualapai Valley (acre-ft/yr)	Sacramento Valley (acre-ft/yr)
1940	<sup>1</sup> <200	<sup>2</sup> <300	<sup>2</sup> <300	1981	<sup>1</sup> <200	(5,000) <sup>6</sup> 6,000	<sup>6</sup> 5,000
1941	<sup>1</sup> <200	<sup>2</sup> <300	<sup>2</sup> <300	1982	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6</sup> 2,000
1942	<sup>1</sup> <200	<sup>2</sup> <300	<sup>2</sup> <300	1983	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6</sup> 2,000
1943	<sup>1</sup> <200	<sup>3</sup> >1,000	<sup>2</sup> <300	1984	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6</sup> 2,000
1944	<sup>1</sup> <200	<sup>3</sup> >1,000	<sup>2</sup> <300	1985	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6</sup> 2,000
1945	<sup>1</sup> <200	<sup>3</sup> >1,000	<sup>4</sup> 250	1986	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6</sup> 2,000
1946	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1987	<sup>1</sup> <200	(6,000) <sup>6</sup> 5,000	<sup>6</sup> 2,000
1947	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1988	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6</sup> 2,000
1948	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1989	<sup>1</sup> <200	<sup>6</sup> 6,000	<sup>6</sup> 2,000
1949	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1990	<sup>1</sup> <200	<sup>6</sup> 6,000	<sup>6</sup> 2,000
1950	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1991	<sup>9</sup> 200	<sup>9</sup> 4,400	<sup>9</sup> 1,800
1951	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1992	<sup>9</sup> 200	<sup>9</sup> 5,400	<sup>9</sup> 2,000
1952	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1993	<sup>9</sup> 200	<sup>9</sup> 5,400	<sup>9</sup> 2,000
1953	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1994	<sup>9</sup> 200	<sup>9</sup> 6,000	<sup>9</sup> 2,200
1954	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1995	<sup>9</sup> 200	<sup>9</sup> 6,300	<sup>9</sup> 2,300
1955	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1996	<sup>9</sup> 200	<sup>9</sup> 7,000	<sup>9</sup> 2,400
1956	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1997	<sup>9</sup> 200	<sup>9</sup> 6,900	<sup>9</sup> 2,200
1957	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1998	<sup>9</sup> 200	<sup>9</sup> 6,800	<sup>9</sup> 2,000
1958	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	1999	<sup>9</sup> 200	<sup>9</sup> 7,500	<sup>9</sup> 2,700
1959	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	2000	<sup>9</sup> 200	<sup>9</sup> 8,200	<sup>9</sup> 2,700
1960	<sup>1</sup> <200	<sup>5</sup> 500	<sup>4</sup> 250	2001	<sup>10</sup> 200	<sup>10</sup> 8,000	<sup>10</sup> 2,800
1961	<sup>1</sup> <200	<sup>4,6</sup> 1,000	<sup>4</sup> 250	2002	<sup>10</sup> 300	<sup>10</sup> 8,500	<sup>10</sup> 2,900
1962	<sup>1</sup> <200	<sup>4,6</sup> 1,000	<sup>4</sup> 250	2003	<sup>10</sup> 300	<sup>10</sup> 8,400	<sup>10</sup> 2,600
1963	<sup>1</sup> <200	<sup>4,6</sup> 1,000	<sup>4</sup> 250	2004	<sup>10</sup> 300	<sup>10</sup> 8,400	<sup>10</sup> 2,900
1964	<sup>1</sup> <200	<sup>4,6</sup> 2,000	<sup>4</sup> 250	2005	<sup>10</sup> 300	<sup>10</sup> 9,100	<sup>10</sup> 2,700
1965	<sup>1</sup> <200	<sup>4,6</sup> 2,000	<sup>4</sup> 4,000	2006	<sup>11</sup> <300	<sup>11</sup> 9,500	<sup>11</sup> 3,600
1966	<sup>1</sup> <200	<sup>4,6</sup> 3,000	<sup>4,7</sup> 4,000	2007	<sup>12</sup> <300	<sup>12</sup> 9,800	<sup>12,13</sup> 4,500
1967	<sup>1</sup> <200	<sup>6,8</sup> 4,000	<sup>6,7,8</sup> 4,000	2008	<sup>12</sup> <300	<sup>12</sup> 9,800	<sup>12,13</sup> 4,500
1968	<sup>1</sup> <200	<sup>6,8</sup> 4,000	<sup>6,7,8</sup> 6,000				
1969	<sup>1</sup> <200	<sup>6,8</sup> 4,000	<sup>6,7,8,5</sup> 5,000				
1970	<sup>1</sup> <200	<sup>6,8</sup> 4,000	<sup>6,7,8,4</sup> 4,000				
1971	<sup>1</sup> <200	<sup>6</sup> 4,000	<sup>6,7,5</sup> 5,000				
1972	<sup>1</sup> <200	<sup>6</sup> 4,000	<sup>6,7,6</sup> 6,000				
1973	<sup>1</sup> <200	<sup>6</sup> 4,000	<sup>6,7,5</sup> 5,000				
1974	<sup>1</sup> <200	<sup>6</sup> 5,000	<sup>6,7,7</sup> 7,000				
1975	<sup>1</sup> <200	<sup>6</sup> 4,000	<sup>6,7,8</sup> 8,000				
1976	<sup>1</sup> <200	<sup>6</sup> 4,000	<sup>6,7,8</sup> 8,000				
1977	<sup>1</sup> <200	<sup>6</sup> 4,000	<sup>6,7,7</sup> 7,000				
1978	<sup>1</sup> <200	<sup>6</sup> 6,000	<sup>6,7,8</sup> 8,000				
1979	<sup>1</sup> <200	<sup>6</sup> 6,000	<sup>6</sup> 6,000				
1980	<sup>1</sup> <200	(6,000) <sup>6</sup> 4,000	<sup>6</sup> 4,000				

<sup>1</sup>Inferred by assuming historic water use less than first-reported values.

<sup>2</sup>Inferred assuming no major water use prior to first-reported values.

<sup>3</sup>Gillespie and Bentley (1971).

<sup>4</sup>Cox and others (1968).

<sup>5</sup>Estimated on the basis of large yield reported for one well in valley.

<sup>6</sup>Anning and Duet (1994).

<sup>7</sup>Pfaff and Clay (1981).

<sup>8</sup>Babcock and others (1969, 1970, 1972).

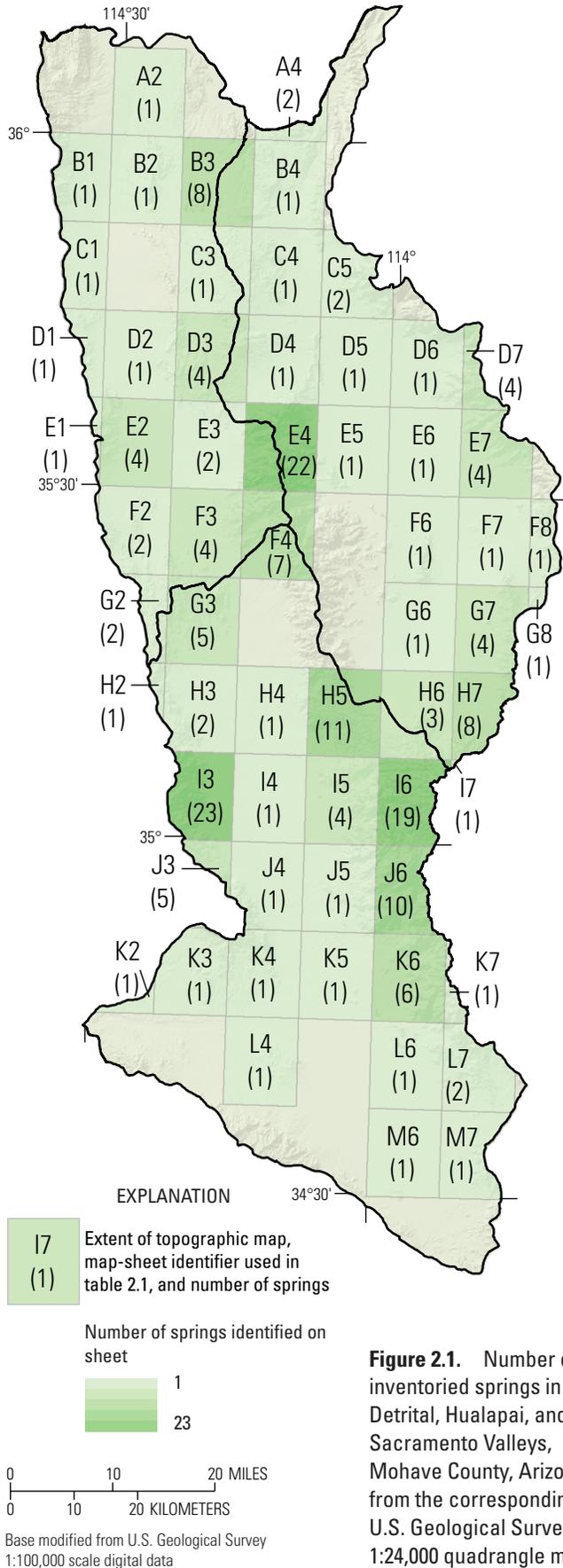
<sup>9</sup>Tadayon (2005).

<sup>10</sup>Compiled using the same methods as Tadayon (2005).

<sup>11</sup>Estimated as the mean of 2005 and 2007–08 values.

<sup>12</sup>Values from this study, reported as mean for 2007–08 values.

<sup>13</sup>Increase caused by inclusion of 1,900 ac-ft/yr industrial water use.



**Figure 2.1.** Number of inventoried springs in Detrital, Hualapai, and Sacramento Valleys, Mohave County, Arizona, from the corresponding U.S. Geological Survey 1:24,000 quadrangle map.

1,000 acre-ft/yr or more, mostly in support of municipal and domestic water uses (Anning and Duet, 1994). By 1991, almost all groundwater pumping withdrawals were for municipal and domestic water uses. Groundwater withdrawals increased steadily from 4,400 acre-ft/yr in 1991 to about 9,800 acre-ft/yr in 2008. Total cumulative groundwater withdrawn from the Hualapai Valley alluvial aquifer from 1940 to 2008 is estimated to be about 270,000 acre-ft.

Sacramento Valley alluvial aquifer groundwater withdrawals were 250 acre-ft/yr or less from 1940 to 1964 (Cox and others, 1968). Groundwater withdrawals increased to 4,000 acre-ft/yr in 1965 and increased again from 1966 to 1978 (4,000–8,000 acre-ft/yr), mostly in support of municipal and domestic water uses for City of Kingman and the surrounding area (Anning and Duet, 1994). Groundwater withdrawals decreased substantially from 1978 to 1982 (from 8,000 to 2,000 acre-ft/yr), likely the result of the City of Kingman limiting pumping of its well field in the volcanic rocks on the east side of Sacramento Valley, owing to depletion of the resource (Rascona, 1991). Since 1982, withdrawals have remained near 2,000 acre-ft/yr, although they have been increasing during that time (Anning and Duet, 1994; Tadayon, 2005). Total cumulative groundwater withdrawn from the Sacramento Valley alluvial aquifer from 1940 to 2008 was about 170,000 acre-ft.

## Appendix 2. Inventory of Springs

Small springs exist in upland areas of the study area because of shallow groundwater discharging either through fractures in mountain-block bedrock or along zones of perched water. A visual inspection of USGS topographic maps (fig. 2.1; table 2.1) was completed to estimate total annual groundwater discharge from upland springs. Field verification of these springs, or an assessment about the completeness of this inventory, is beyond the scope of this report. Springs were excluded if their elevation and surrounding terrain indicated they most likely discharge Lake Mead bank-storage water or alluvial-aquifer groundwater. Within the study area, 204 springs were identified, of which 19 percent, 29 percent, and 52 percent were in Detrital, Hualapai, and Sacramento Valleys, respectively.

Median upland-spring discharge for each valley was estimated by using spring-discharge data from the Arizona Department of Water Resources (2009b). Four springs in Detrital Valley had a median discharge of 4.5 gal/min. Seventeen springs in Hualapai Valley had a median discharge of 3.0 gal/min. Fifty-eight springs in Sacramento Valley had a median discharge of 3.0 gpm. Multiplying these values by the number of inventoried springs in each valley resulted in upland-spring discharge estimates of 300 acre-ft/yr for Detrital Valley, 300 acre-ft/yr for Hualapai Valley, and 500 acre-ft/yr for Sacramento Valley. These values are low compared to total recharge and runoff in the study area (about 4 percent).

Several springs have been identified in the vicinity of Lake Mead (table 2.2). These are discussed in the “Springs Near Lake Mead” section. They are divided into springs that discharge above and below the maximum possible surface-water elevation of Lake Mead.

**Table 2.1. Inventory of upland springs, identified by visual inspection of U.S. Geological Survey 1:24,000 topographic map sheets, Mohave County, Arizona. Names provided where available, otherwise general location indicated.**

[\*\*, multiple springs within the same vicinity; DV, Detrital valley; HV, Hualapai Valley; SV, Sacramento Valley; NW, northwest; NE, northeast; SW, southwest; SE, southeast]

Valley	Map sheet identifier <sup>1</sup>	Topographic map name	Spring name (number of springs, if more than 1)	Valley	Map sheet identifier <sup>1</sup>	Topographic map name	Spring name (number of springs, if more than 1)
DV	A2	Bonelli Bay	Alkali Spring	HV	E4	Mount Tipon	** Unnamed, near Cane Spring (2)
DV	B1	Mount Wilson	Unnamed	HV	E4	Mount Tipon	Cliff Spring
DV	B2	Senator Mountain NW	Unnamed	HV	E4	Mount Tipon	Tompson Spring
DV	B3	Senator Mountain NE	Salt Spring	HV	E4	Mount Tipon	Canyon Spring
DV	B3	Senator Mountain NE	** Unnamed, near Salt Spring (7)	HV	E4	Mount Tipon	Cottonwood Spring
DV	C1	Housholder Pass	Unnamed	HV	E4	Mount Tipon	Unnamed, near Cottonwood Spring
DV	D1	Mohave Mine	Unnamed	HV	E5	Mount Tipon SE	Unnamed
DV	D2	White Hills West	Unnamed	HV	E6	Music Mountains SW	Black Rock Spring
DV	D3	White Hills East	Alta Spring	HV	E7	Music Mountains SE	Cedar Spring
DV	D3	White Hills East	Mountain Spring	HV	E7	Music Mountains SE	Pole Corral Spring
DV	E1	Mount Perkins	Unnamed	HV	E7	Music Mountains SE	Kohinoor Spring
DV	E2	Middle Water Spring	Kemple Spring	HV	E7	Music Mountains SE	Coyote Spring
DV	E2	Middle Water Spring	Middle Water Spring	HV	F4	Chloride	Fall spring
DV	E2	Middle Water Spring	** Unnamed, Middle Water Spring (2)	HV	F4	Chloride	Cedar Spring
DV	E3	Dolan Springs	Sacaton Spring	HV	F4	Chloride	Vock Spring
DV	E3	Dolan Springs	Mud Spring	HV	F4	Chloride	Unnamed, near Lewis Ranch
DV	E4	Mount Tipon	** Antelope Springs (3)	HV	F4	Chloride	Gaddis Spring
DV	E4	Mount Tipon	** Dolan Springs (3)	HV	F6	Long Mountain	Unnamed
DV	E4	Mount Tipon	Upper Indian Spring	HV	F7	Antares	Unnamed
DV	E4	Mount Tipon	Lower Indian Spring	HV	F8	Valentine	Stone Corral Spring
DV	E4	Mount Tipon	Arizona Spring	HV	G6	City of Kingman Airport	Unnamed
DV	E4	Mount Tipon	Unnamed, near Arizona Spring	HV	G7	Peacock Peak	Crecent Spring
DV	F2	Grasshopper Junction NW	Unnamed	HV	G7	Peacock Peak	North Water Spring
DV	F3	Grasshopper Junction	Quail Spring	HV	G7	Peacock Peak	Willow Spring
DV	F3	Grasshopper Junction	** Unnamed, near Quail Spring (2)	HV	G7	Peacock Peak	Mud Spring
DV	F4	Chloride	Putman spring	HV	G8	Hackberry	Unnamed
HV	A4	Hiller Mountains	Burro Spring	HV	H6	Rattlesnake Hill	Falls Spring
HV	A4	Hiller Mountains	Hualapai Wash Spring	HV	H6	Rattlesnake Hill	Unnamed, near Falls Spring
HV	B4	Garnet Mountain NW	Unnamed	HV	H7	Hualapai Spring (Peak NE)	Hualapai Spring
HV	C3	Senator Mountain	Eileen Spring	HV	H7	Hualapai Spring (Peak NE)	Burn Spring
HV	C4	Gold Basin	Unnamed	HV	H7	Hualapai Spring (Peak NE)	** Unnamed, near Burn Spring (2)
HV	C5	Garnet Mountain	Rattlesnake Spring	HV	H7	Hualapai Spring (Peak NE)	Lion Kill Spring
HV	C5	Garnet Mountain	Valley Ranch Spring	HV	H7	Hualapai Spring (Peak NE)	** Unnamed, near Lion Kill Spring (3)
HV	D3	White Hills East	Red Willow Spring	HV	I7	Dean Peak	Yellow Pine Spring
HV	D3	White Hills East	Rock Spring	SV	F2	Grasshopper Junction NW	Unnamed
HV	D4	Mt. Tipon NW	Unnamed	SV	F3	Grasshopper Junction	Unnamed
HV	D5	Red Lake	Unnamed	SV	F4	Chloride	Swicher Spring
HV	D6	Music Mountains NW	Fred Jackson Spring	SV	G2	Burns Spring	Willow Spring
HV	D7	Music Mountains NE	** Clay Springs (4)	SV	G2	Burns Spring	Burro Spring
HV	E4	Mount Tipon	Grapvine Spring	SV	G3	Grasshopper Junction SE	Mud Spring
HV	E4	Mount Tipon	Ash Creek Spring				
HV	E4	Mount Tipon	** Unnamed, near Ash Creek Spring (2)				
HV	E4	Mount Tipon	Cane Spring				

**Table 2.1. Inventory of upland springs, identified by visual inspection of U.S. Geological Survey 1:24,000 topographic map sheets, Mohave County, Arizona. Names provided where available, otherwise general location indicated.—Continued**

[\*\*, multiple springs within the same vicinity; DV, Detrital valley; HV, Hualapai Valley; SV, Sacramento Valley; NW, northwest; NE, northeast; SW, southwest; SE, southeast]

Valley	Map sheet identifier <sup>1</sup>	Topographic map name	Spring name (number of springs, if more than 1)	Valley	Map sheet identifier <sup>1</sup>	Topographic map name	Spring name (number of springs, if more than 1)
SV	G3	Grasshopper Junction SE	Cottonwood Spring	SV	I6	Hualapai Peak	** Unnamed, near Yellowflower Creek (5)
SV	G3	Grasshopper Junction SE	Thorne Spring	SV	I6	Hualapai Peak	Yellowflower Spring
SV	G3	Grasshopper Junction SE	Little Thorne Spring	SV	I6	Hualapai Peak	Cottonwood Spring
SV	G3	Grasshopper Junction SE	Red Rock Spring	SV	I6	Hualapai Peak	Grande Spring
SV	H2	Union Pass	Unnamed, at Union Pass	SV	I6	Hualapai Peak	Black Inky Spring
SV	H3	Secret Pass	Cave Spring	SV	I6	Hualapai Peak	Willow Spring
SV	H3	Secret Pass	Trough Spring	SV	I6	Hualapai Peak	Unnamed, near Walnut Creek
SV	H4	Kingman NW	Jones Spring	SV	I6	Hualapai Peak	Walnut Creek Mine Spring
SV	H5	Kingman	Oak Spring	SV	I6	Hualapai Peak	** Unnamed, in Whiskey Basin (2)
SV	H5	Kingman	Gross Spring	SV	J3	Warm Springs	Antelope Spring
SV	H5	Kingman	Unnamed, west of Gross Spring	SV	J3	Warm Springs	Alkali Spring
SV	H5	Kingman	Coyote Holes Spring	SV	J3	Warm Springs	** Warm Springs (3)
SV	H5	Kingman	Johnson Spring	SV	J4	Yucca NW	Caliche Spring
SV	H5	Kingman	Unnamed, south of Johnson Spring	SV	J5	Yucca NE	Unnamed
SV	H5	Kingman	Grapevine Spring	SV	J6	Wabayuma Peak	Unnamed, north of Wabayuma Peak
SV	H5	Kingman	Atlantic Spring	SV	J6	Wabayuma Peak	Grapevine Spring
SV	H5	Kingman	Unnamed, in Cook Canyon north of I-40	SV	J6	Wabayuma Peak	Cottonwood Spring
SV	H5	Kingman	** Unnamed, east of Harris south of I-40 (2)	SV	J6	Wabayuma Peak	Unnamed, near Boriana Mine
SV	H6	Rattlesnake Hill	Shack Spring	SV	J6	Wabayuma Peak	** Unnamed, near Mackenzie Creek (2)
SV	I3	Mount Nutt	Baker Spring	SV	J6	Wabayuma Peak	** Bee Springs (2)
SV	I3	Mount Nutt	** Walker Springs (2)	SV	J6	Wabayuma Peak	Unnamed, near Copper Creek
SV	I3	Mount Nutt	Cool Spring	SV	J6	Wabayuma Peak	Bateman Spring
SV	I3	Mount Nutt	** Unnamed, in Black Mountains (3)	SV	K2	Warm Springs SW	Unnamed
SV	I3	Mount Nutt	Cave Spring	SV	K3	Warm Springs SE	Unnamed
SV	I3	Mount Nutt	** Fig Springs (2)	SV	K4	Yucca	Unnamed
SV	I3	Mount Nutt	Dripping Spring	SV	K5	Yucca SE	Unnamed
SV	I3	Mount Nutt	Gold Trail Spring	SV	K6	Creamery Canyon	Blue Mary Spring
SV	I3	Mount Nutt	** Twin Springs (2)	SV	K6	Creamery Canyon	Unnamed, north of Blue Mary Spring
SV	I3	Mount Nutt	** Indian Springs (2)	SV	K6	Creamery Canyon	Cottonwood Spring
SV	I3	Mount Nutt	Shaffer Spring	SV	K6	Creamery Canyon	Unnamed, near Round Valley
SV	I3	Mount Nutt	** Unnamed, east of Shaffer Spring (3)	SV	K6	Creamery Canyon	** Unnamed, near Cow Creek (2)
SV	I3	Mount Nutt	** Unnamed, near Sitgreaves Pass (3)	SV	K7	Diamond Joe Peak	Grapevine Spring
SV	I4	Kingman SW	Unnamed	SV	L4	Buck Mountains	Unnamed
SV	I5	Kingman SE	Lookout Spring	SV	L6	Dutch Flats NW	Unnamed
SV	I5	Kingman SE	** Unnamed, near south Walnut Creek (3)	SV	L7	Beecher Canyon	Mud Spring
SV	I6	Hualapai Peak	Wiskey Spring	SV	L7	Beecher Canyon	Beecher Spring
SV	I6	Hualapai Peak	Unnamed, above Camp Stephens	SV	M6	Dutch Flats SW	Unnamed
SV	I6	Hualapai Peak	** Unnamed, above Camp Levi (3)	SV	M7	Dutch Flats SE	Unnamed

<sup>1</sup>Corresponds to identifiers shown on Figure 2.1.

**Table 2.2.** Inventory of springs near the southern shoreline of Lake Mead, Mohave County, Arizona; includes measured discharge, where available.

[discharge data 1960–80 from Laney (1979); ?, unknown month for discharge value]

Spring Name	Elevation, in feet above mean sea level	Rock type from which spring discharges	Latitude and longitude (North American Datum, 1927)	Discharge (month/year: gallons per minute)
<b>Above Lake Mead maximum lake-surface elevation</b>				
Burro Spring	1,360	Granite	36°00'47" N 114°09'51" W	6/73: 5; 2/08: 38; 5/09: 16
Salt Spring	1,390	Alluvium	35°59'19" N 114°15'19" W	5/73: 9; 5/75: 2; 2/08: 8; 5/09: 0
unnamed	3,230	Granite	35°59'27" N 114°35'06" W	2/72: 0.3
Alkali Spring	1,220	Muddy Creek Formation	36°02'54" N 114°27'21" W	4/08: 0
<b>Below Lake Mead maximum lake-surface elevation</b>				
Hualapai Wash Seeps	1,214	Muddy Creek Formation	36°00'43" N 114°07'36" W	4/08: 0
Granite Cove	1,197	Granite	36°00'42" N 114°08'56" W	2/76: 25; 4/08: 0
Teal Cove	1,180	Volcanics	36°04'24" N 114°20'43" W	4/08: 0.2
Monkey Cove	1,150	Alluvium	36°02'30" N 114°19'54" W	?/64: 1,200; 2/08: 1,350; 5/09: 380

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