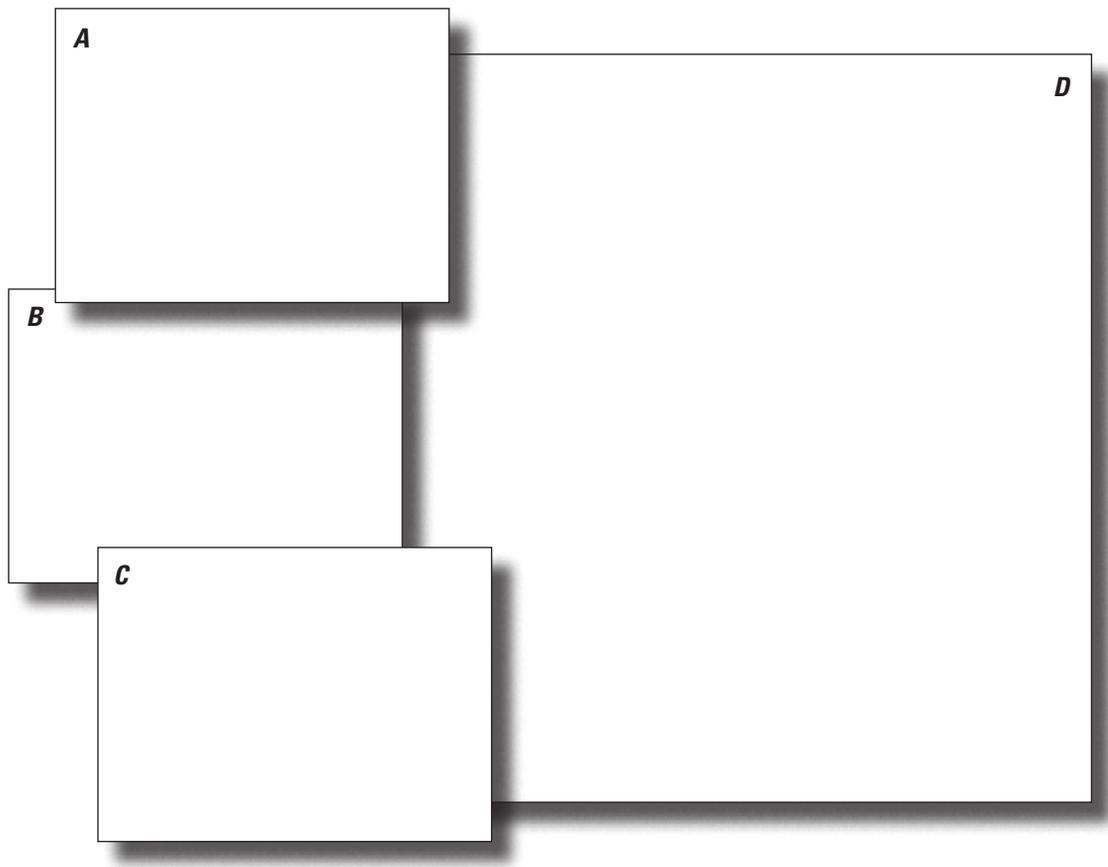


Prepared in cooperation with the United States Fish and Wildlife Service

Physical Characteristics of the Lower San Joaquin River, California, in Relation to White Sturgeon Spawning Habitat, 2011–14



Scientific Investigations Report 2017–5069



Cover. Front cover photographs taken by Mathieu Marineau: *A*, Gravel bar in the San Joaquin River at the mouth of Orestimba Creek. *B*, Close up view of cobble and gravel found in the bed of Del Puerto Creek, near the confluence with the San Joaquin River. *C*, Bed material sampler during field work in the San Joaquin River. *D*, Oblique aerial image showing the San Joaquin River at Sturgeon Bend and confluence of the Stanislaus River (NAIP, 2012).

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By Mathieu D. Marineau, Scott A. Wright, Daniel R. Whealdon-Haught, and Paul J. Kinzel

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U.S. Department of the Interior
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Contents

Abstract.....	1
Introduction.....	1
Description of the San Joaquin River Basin	2
Lower San Joaquin River and its Tributaries	2
Study Area, River Reaches, and Subreaches	5
Hydrology	8
White Sturgeon Ecology.....	11
White Sturgeon Monitoring in the San Joaquin River by the U.S. Fish and Wildlife Service	12
Study Methods	12
Vertical and Horizontal Datum	12
River-Centerline Convention	12
Equipment.....	14
Multibeam Echosounder System	14
Acoustic Doppler Current Profiler	14
Side-Scanning Sonar	14
Data Collection.....	15
Water Depth and Velocity Measurements using Acoustic Doppler Current Profiler, 2011–14.....	15
Bathymetry Mapping using Multibeam Sonar, 2011	17
Bed-Material Mapping using Acoustic Backscatter in 2012 and Side-Scan Sonar in 2013.....	17
Tributary Longitudinal Elevation and Slope Profiles	17
Bed-Material Sampling, 2012–14	19
Bed-Material Sample Preparation and Particle-Size Analysis.....	21
Summary of Channel Characteristics, Stream Slope, and Bed Aggradation in the Lower San Joaquin River.....	22
Depth and Velocity.....	22
Water Year 2011: Wet.....	22
Water Years 2012–14: Dry and Critical	22
Bed-Material Composition and Sources	22
Stream Slope	35
Bed Aggradation, 2011–14.....	35
Discussion and Conclusions: Bed Material and Streamflow in the Lower San Joaquin River	38
Gravel in a Sand-Bedded River	38
Changes in Hydrology	38
Aggradation of the River Bed	38
Sources of Gravel and Cobble.....	38
Summary.....	39
References Cited.....	40
Appendix 1: River Kilometer Station Locations.....	45

Figures

1. Map showing the San Joaquin Basin, Tulare Basin, and lower San Joaquin River study area, California	3
2. Photograph showing large, broken concrete slabs, which are used as bank protection in many areas of the lower San Joaquin River, California	4
3. Photograph showing stones (mostly cobble size), which are used as bank protection in many areas of the lower San Joaquin River, California	4
4. Photograph showing typical consolidated bank on the lower San Joaquin River, California	5
5. Map showing lower San Joaquin River reaches, subreaches of interest, locations of selected U.S. Geological Survey and California Department of Water Resources streamgages, river kilometer stations, and 2011–12 U.S. Fish and Wildlife Service sturgeon-egg sampling sites	6
6. Hydrographs showing annual peak streamflow, measured as instantaneous discharge at selected streamgages, and the time of dam completion, where applicable, in the San Joaquin Basin, California, 1920–2010	9
7. Hydrographs showing daily average streamflow, measured as discharge at the U.S. Geological Survey streamgage at Vernalis (station 11303500), for the lower San Joaquin River, California, categorized by water-year type for water years 1924, 1930–67 and 1968–2014	10
8. Hydrographs showing the daily mean streamflow, measured as discharge, during the study (water years 2011–14) relative to the historical range of discharge measured on the San Joaquin River, California, at the U.S. Geological Survey streamgage at Vernalis (station 11303500)	11
9. Hydrograph and maps showing the dates and locations of the U.S. Fish and Wildlife Service sampling and U.S. Geological Survey data collection, 2011–12	13
10. Maps showing locations where water depth and velocity data were collected, by year, in the lower San Joaquin River, California, 2011–14	16
11. Images showing examples of side-scan sonar from a sand-bedded section in the Grayson Bridge subreach of the lower San Joaquin River, California, April 23, 2013	18
12. Image showing plan view of side-scan sonar of a gravel-bedded section in the Grayson Bridge subreach of the lower San Joaquin River, California, April 23, 2013	19
13. Maps showing sites where bed-material samples were collected in the lower San Joaquin River, Del Puerto Creek, and Orestimba Creek, California	20
14. Maps showing water depth in the lower San Joaquin River, California	23
15. Maps showing depth-averaged velocity in the lower San Joaquin River, California	24
16. Images showing water depth and bed-material size class in the Sturgeon Bend subreach of the lower San Joaquin River, California	25
17. Images showing water depth and bed-material size class in the Old Fisherman’s Club subreach of the lower San Joaquin River, California	26
18. Images showing water depth and bed-material size class in the Grayson Bridge subreach of the lower San Joaquin River, California	27
19. Images showing streamflow depth-averaged velocity and bed-material size class in the Sturgeon Bend subreach of the lower San Joaquin River, California	28

Figures—Continued

20.	Images showing streamflow depth-averaged velocity and bed-material size class in the Old Fisherman’s Club subreach of the lower San Joaquin River, California.....	29
21.	Images showing streamflow depth-averaged velocity and bed-material size class in the Grayson Bridge subreach of the lower San Joaquin River, California	30
22.	Photograph showing cobble in the dry bed of Orestimba Creek at Eastin Road, approximately 5 kilometers northwest of Newman, California, April 24, 2014	31
23.	Photograph showing the edge of a large gravel bar in the lower San Joaquin River at river kilometer 146.3, the mouth of Del Puerto Creek, California, November 8, 2013.....	31
24.	Photograph showing the mouth of Del Puerto Creek, California, November 8, 2013, where the channel cut through a large gravel deposit.....	32
25.	Photograph showing the dry channel bed of Del Puerto Creek, 22 kilometers upstream from its confluence with the lower San Joaquin River, California, November 1, 2013.....	32
26.	Photograph showing Del Puerto Creek at Rogers Road, November 1, 2013	33
27.	Photograph showing an exposed gravel bar in the lower San Joaquin River, California, near river kilometer 171.8, the mouth of Orestimba Creek, May 13, 2014	33
28.	Graphs summarizing particle-size distribution of bed material samples collected from channels during 2011–13.....	34
29.	Graphs showing longitudinal profiles of water-surface elevation, water-surface slope, and principal dams in rivers and creeks of the San Joaquin Basin, California....	36
30.	Graph showing bed-elevation changes along selected longitudinal section in the Sturgeon Bend subreach, lower San Joaquin River, California, 2011–14	37
31.	Graph showing bed-elevation changes along selected longitudinal section in the Grayson Bridge subreach, San Joaquin River, California, 2011–14	37

Tables

1.	River reaches and subreaches of the lower San Joaquin River and the respective downstream and upstream river kilometer stations.....	7
2.	Selected landmarks and geographical features along the lower San Joaquin River, by river kilometers and geographical coordinates	7
3.	Summary of hydroacoustic data collection sites, including daily mean streamflow, measured as discharge, in the reach during data collection and the exceedance probability of that discharge	15
4.	Description and size range of sediment particles in particle-size categories used in this report.....	21
1–1.	River kilometer stations, and their associated geographic coordinates, based on the 2012 National Aerial Imagery Program aerial imagery of the lower San Joaquin River	45

Conversion Factors

International System of Units to U.S. customary units

Multiply	By	To obtain
Length		
centimeter (cm)	0.3937	inch (in.)
millimeter (mm)	0.03937	inch (in.)
meter (m)	3.281	foot (ft)
kilometer (km)	0.6214	mile (mi)
meter (m)	1.094	yard (yd)
Area		
square kilometer (km ²)	247.1	acre
square kilometer (km ²)	0.3861	square mile (mi ²)
Volume		
cubic meter (m ³)	35.31	cubic foot (ft ³)
cubic meter (m ³)	1.308	cubic yard (yd ³)
cubic meter (m ³)	0.0008107	acre-foot (acre-ft)
cubic kilometer (km ³)	810,714	acre-foot (acre-ft)
cubic kilometer (km ³)	0.2399	cubic mile (mi ³)
Flow rate		
cubic meter per second (m ³ /s)	70.07	acre-foot per day (acre-ft/d)
meter per second (m/s)	3.281	foot per second (ft/s)
cubic meter per second (m ³ /s)	35.31	cubic foot per second (ft ³ /s)
millimeter per year (mm/yr)	0.03937	inch per year (in/yr)

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as

$$^{\circ}\text{F} = (1.8 \times ^{\circ}\text{C}) + 32.$$

Datum

Vertical coordinate information is referenced to the North American Vertical Datum of 1988 (NAVD 88).

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Elevation, as used in this report, refers to distance above the vertical datum.

Abbreviations

ADCP	acoustic Doppler current profiler
GNSS	Global Navigation Satellite System
GPS	global positioning system
kHz	kilohertz
MAF	million acre feet
MHz	megahertz
MRU	motion reference unit
NAIP	National Aerial Imagery Program
Rkm	river kilometer
RO	annual runoff
RTK-GPS	real-time kinematic global positioning system
RTN-GPS	real-time network global positioning system
USFWS	U.S. Fish and Wildlife Service
USGS	U.S. Geological Survey
WY	water year

Physical Characteristics of the Lower San Joaquin River, California, in Relation to White Sturgeon Spawning Habitat, 2011–14

By Mathieu D. Marineau, Scott A. Wright, Daniel R. Whealdon-Haught, and Paul J. Kinzel

Abstract

The U.S. Fish and Wildlife Service confirmed that white sturgeon (*Acipenser transmontanus*) recently spawned in the lower San Joaquin River, California. Decreases in the San Francisco Bay estuary white sturgeon population have led to an increased effort to understand their migration behavior and habitat preferences. The preferred spawning habitat of other white sturgeon (for example, those in the Columbia and Klamath Rivers) is thought to be areas that have high water velocity, deep pools, and coarse bed material. Coarse bed material (pebbles and cobbles), in particular, is important for the survival of white sturgeon eggs and larvae. Knowledge of the physical characteristics of the lower San Joaquin River can be used to preserve sturgeon spawning habitat and lead to management decisions that could help increase the San Francisco Bay estuary white sturgeon population.

Between 2011 and 2014, the U.S. Geological Survey, in cooperation with the U.S. Fish and Wildlife Service, assessed selected reaches and tributaries of the lower river in relation to sturgeon spawning habitat by (1) describing selected spawning reaches in terms of habitat-related physical characteristics (such as water depth and velocity, channel slope, and bed material) of the lower San Joaquin River between its confluences with the Stanislaus and Merced Rivers, (2) describing variations in these physical characteristics during wet and dry years, and (3) identifying potential reasons for these variations.

The lower San Joaquin River was divided into five study reaches. Although data were collected from all study reaches, three subreaches where the USFWS collected viable eggs at multiple sites in 2011–12 from Orestimba Creek to Sturgeon Bend were of special interest. Water depth and velocity were measured using two different approaches—channel cross sections and longitudinal profiles—and data were collected using an acoustic Doppler current profiler.

During the first year of data collection (water year 2011), runoff was greatest, and gaged streamflow, measured as discharge, peaked at 875 cubic meters per second in the lower San Joaquin River. Also during that year, water velocity was generally between 0.6 and 0.9 meters per second, and depth was typically between 2.5 and 4.5 meters, but water depth

exceeded 6 meters in several pools. Water year 2011 was classified as a “wet” year. Later water years were classified as either “dry” (water year 2012) or “critical” (water years 2013 and 2014). During the drier years, water was shallower, and velocities were slower. The streambed aggraded in several areas during the study. At Sturgeon Bend, for example, which had the deepest pool measured in 2011 (maximum depth was 14 meters), about 8 meters of sediment was deposited by 2014.

The bed of the lower San Joaquin River was predominately sand, except in areas downstream from the mouth of Del Puerto Creek. A large amount of sand, gravel, and cobble was deposited at the mouth of Del Puerto Creek, and in the 9.5 kilometers downstream from the mouth of Del Puerto Creek, we encountered several gravel bars and patches of gravel-size (8–64 millimeters) bed material. Del Puerto and Orestimba Creeks drain from the Coast Ranges on the west side of the river. Only small quantities of gravel-size bed material were observed in the reach downstream from Orestimba Creek, indicating Orestimba Creek does not deliver much coarse sediment to the lower San Joaquin River. Del Puerto Creek appeared to be the primary source of gravels suitable for white sturgeon spawning in the lower San Joaquin River, and thus, it is important for the long-term spawning success of sturgeon in the San Joaquin River.

Introduction

White sturgeon (*Acipenser transmontanus*) have long been known to inhabit the San Francisco Bay estuary and migrate up the Sacramento River to spawn (Kohlhorst, 1976; Schaffter, 1997). They have been known to migrate up the lower San Joaquin River (Moore and others, 1990), but it was only speculated that they actually spawn there (for example, Kohlhorst, 1976; Schaffter, 1997). Viable white sturgeon eggs were collected by the U.S. Fish and Wildlife Service (USFWS) at one sampling site in 2011 and at four sampling sites in 2012, providing evidence of spawning activity in the lower San Joaquin River (Gruber and others, 2012; Jackson and Van Eenennaam, 2013).

White sturgeon face many stressors, including fishing, poaching, degraded habitat, and recruitment failures (Birstein, 1993; Gross and others, 2002; Jackson and others, 2016). Decreases in the population of white sturgeon in the Sacramento–San Joaquin drainage (Central Valley watershed) have been a management concern for the last century (Klimley and others, 2015).

Given the evidence of spawning adults, managers are interested in learning more about the physical characteristics of the lower San Joaquin River, particularly the hydraulic conditions and bed-material sediment types. Research in the Columbia and Kootenai Rivers of the Pacific Northwest indicated that high streamflow velocity and coarse bed material increased the survival of white sturgeon eggs and larvae (Parsley and others, 1993; Parsley and Beckman, 1994; Billard and Lecointre, 2000). Identifying the physical characteristics of the river where white sturgeon spawn can help protect their habitat and increase their survival to the larval drift stage of the life cycle. In response to the need to characterize the physical habitat and bed-material sediment characteristics of the lower San Joaquin River, the U.S. Geological Survey (USGS) initiated this study in cooperation with the U.S. Fish and Wildlife Service (USFWS). This study investigated physical characteristics of selected reaches of the lower San Joaquin River in relation to sturgeon spawning habitat.

The purposes of this report are (1) to describe select spawning sites in terms of habitat-related physical characteristics (such as stream depth, velocity, and bed material) in the lower San Joaquin River between the confluences of the Stanislaus and Merced Rivers, (2) to document variations in these physical characteristics during wet and dry years, and (3) to identify potential factors linked to these variations.

The physical characteristic data also can be used to develop habitat-suitability curves for white sturgeon in the lower San Joaquin River and to help develop hydraulic and sediment transport models for reaches of interest. In this report, we describe hydraulic conditions, bed-material characteristics, and the hydrologic history of the lower San Joaquin River in the context of spawning preferences of white sturgeon. Neither the physical transport models nor habitat-suitability curves were developed as part of this report, however.

Description of the San Joaquin River Basin

The San Joaquin Valley is in the southern part of the Central Valley watershed (fig. 1). The San Joaquin Valley can be divided into two large basins: the San Joaquin Basin to the north and the Tulare Basin to the south. Historically, the Tulare Basin had a large waterbody known as Tulare Lake, into which streams flowed that drained the Tulare Basin. The water level in the lake fluctuated depending on inflows from precipitation and snowmelt runoff, and water would often

overflow to Fresno Slough, toward the San Joaquin River (Tinkham, 1921). Agricultural diversions of Tulare Lake's tributary rivers and streams, however, caused the lake to dry out by the end of the 19th century (Wright, 1899). In addition, the combination of dam construction, surface-water diversion, and groundwater extraction resulted in the loss of hydrologic connectivity between the Tulare Basin and the San Joaquin Basin, except when the King River occasionally floods, which can still contribute streamflow to the San Joaquin Basin.

Lower San Joaquin River and its Tributaries

The lower San Joaquin River begins around river kilometer (Rkm) 234 and flows to the Sacramento–San Joaquin Delta to its confluence with the Sacramento River at Rkm 0. Upstream from Rkm 234, the San Joaquin River historically flowed through a low-lying area and had multiple interconnected channels that frequently overflowed to the surrounding land, forming freshwater wetlands. Upstream from this section, flow in the San Joaquin River is regulated by Friant Dam, and much of the water released is diverted at the Mendota Dam and Sack Dam for agricultural use (California Regional Water Quality Control Board, 1989). During high flows, water is diverted at the San Joaquin River Control Structure (California Department of Water Resources, 2010), but the diverted water eventually flows back to the San Joaquin River through the Mariposa Bypass (Rkm 234).

The lower San Joaquin River is a low-gradient, single-channel, generally sand-bedded, meandering river. Although most of the banks are natural, there are large sections that have revetted banks (that is, sloping banks covered with large rocks to reduce bank erosion and river migration; figs. 2, 3). Sections where both banks are revetted tend to be narrower than sections with natural banks. In areas of active channel migration, mature trees often fall in the river channel and can remain for years (fig. 4).

Streamflow in the lower San Joaquin River is primarily from the Stanislaus, Tuolumne, and Merced Rivers; from the “middle” San Joaquin River through the Mariposa Bypass; and, to a lesser degree, from Del Puerto and Orestimba Creeks. The four rivers flow west from the Sierra Nevada, and all are regulated by dams. By contrast, the two creeks flow east from the Coast Ranges, do not have any dams, and are ephemeral. Streamflow in the lower San Joaquin River is also affected by numerous agricultural diversions (California Regional Water Quality Control Board, 1989).

There are a few other streams that flow from the Coast Ranges (for example, Hospital and Ingram Creeks), which drain smaller watersheds and appear to have more poorly defined channels in the flatter areas of the valley floor. Most of these creeks have been diverted into agricultural water-supply canals. Early land maps and historical accounts do not show or mention Hospital and Ingram Creeks (formerly *Arroyo de Ospital* and *Arroyo de la Suerte*, respectively) as tributaries of the San Joaquin River (for example, Grimes, 1858; Hall, 1886; Reed and others, 1891).



Base modified from U.S. Geological Survey and other Federal and State digital data, various scales; Albers Equal Area Conic projection, standard parallels are 29°30" and 45°30"; North American Datum of 1983

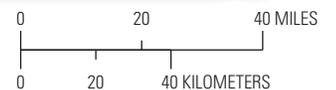


Figure 1. San Joaquin Basin, Tulare Basin, and lower San Joaquin River study area, California.



Figure 2. Large, broken concrete slabs, which are used as bank protection in many areas of the lower San Joaquin River, California.



Figure 3. Stones (mostly cobble size), which are used as bank protection in many areas of the lower San Joaquin River, California.



Figure 4. Typical consolidated bank on the lower San Joaquin River, California. The top of the bank in this photo is approximately 2–3 meters above the water surface. A partially submerged, fallen tree is in the foreground.

Study Area, River Reaches, and Subreaches

The lower San Joaquin River study area extends from the mouth of the Merced River downstream to the USGS streamgage at Vernalis (station 11303500, San Joaquin River near Vernalis, CA), which is on the lower San Joaquin River, downstream from the mouth of the Stanislaus River. The lower San Joaquin River was divided into five reaches for this report. These reach divisions were based on the tributary confluences. The Tuolumne Reach, for example, is the reach of the lower San Joaquin River between the mouths of the Tuolumne River (Rkm 130.6) and the Stanislaus River (Rkm 116.3). The five reaches are shown on the study-area map (fig. 5) and are listed, with their extent in river kilometers, in table 1.

During the 4 years of this study, data were collected throughout the study area, focusing on the reaches with USFWS sampling sites. The USFWS sampled several areas between Rkm 115.2 and 145.3 during 2011 and 2012 (Gruber and others, 2012; Jackson and Van Eenennaam, 2013); the sampling sites were based on pool habitat, water velocity, or previous observations of sturgeon (Jackson and others, 2016). The USFWS collected viable eggs at five sampling sites (Jackson and others, 2016), which can be loosely grouped into

three subreaches of the river. The USGS collected data for 3 to 4 years in those three subreaches, whereas in other areas, data were often only collected for 1 or 2 years. The three, primary study-area subreaches are referenced throughout this report, and the locations of these three subreaches (in river kilometers) are listed in table 1.

The subreach sampled farthest downstream is near Sturgeon Bend. The Sturgeon Bend subreach is the section of the lower San Joaquin River generally between the confluence of the Stanislaus River and about 2 km upstream from the Airport Way Bridge (Rkm 116.3–114.0). Sturgeon Bend is a sharp bend in the lower San Joaquin River, which is popular with local fisherman. The second-most downstream subreach is the Old Fisherman’s Club subreach, between Finnegan Cut (Rkm 127.5) and the upstream end of the straight channel next to Highway 132 (Rkm 125.2), next to the Old Fisherman’s Club, a private sportsman’s club on the east bank of the lower San Joaquin River. The third subreach is near Grayson Bridge between the Grayson Road Bridge (Rkm 139.9) and Rkm 137.3. The locations (in river kilometers) of other selected landmarks and geographical features are listed in table 2.

6 Physical Characteristics of the Lower San Joaquin River, California, in Relation to White Sturgeon Spawning Habitat, 2011–14

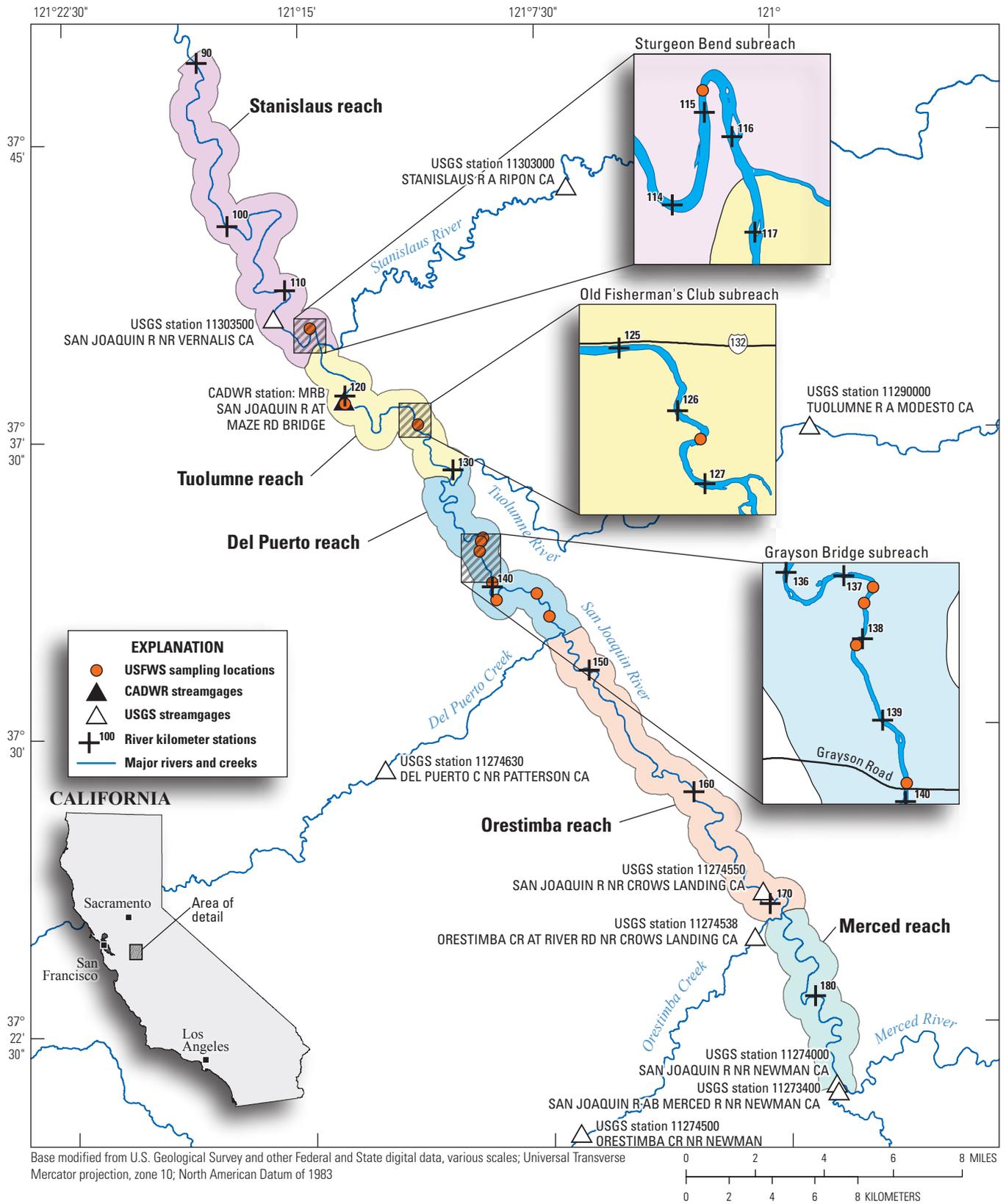


Figure 5. Lower San Joaquin River reaches, subreaches of interest (shown in three upper-right inset maps), locations of selected U.S. Geological Survey (USGS) and California Department of Water Resources (CADWR) streamgages, river kilometer stations, and 2011–12 U.S. Fish and Wildlife Service (USFWS) sturgeon-egg sampling sites.

Table 1. River reaches and subreaches of the lower San Joaquin River and the respective downstream and upstream river kilometer stations.

[River station based on channel configuration during 2012 National Aerial Imagery Program aerial imagery.]

Reach or subreach name	River kilometer	
	Downstream	Upstream
Stanislaus reach	85.1	116.3
Tuolumne reach	116.3	130.6
Del Puerto reach	130.6	146.3
Orestimba reach	146.3	171.8
Merced reach	171.8	187.6
Sturgeon Bend subreach	114.0	116.3
Old Fisherman's Club subreach	125.2	127.5
Grayson Bridge subreach	137.3	139.9

Table 2. Selected landmarks and geographical features along the lower San Joaquin River, by river kilometers and geographical coordinates.

[River stationing based on channel configuration during 2012 National Aerial Imagery Program aerial imagery. North American Datum of 1983. Easting and northing are map grid coordinates based on Universal Transverse Mercator projection, zone 10N. **Abbreviations:** dd, decimal degrees; Rkm, river kilometer]

Feature	Rkm	Latitude (dd)	Longitude (dd)	Northing (meters)	Easting (meters)
Confluence with Sacramento River	0.2	38.065	-121.854	4213606	600512
Bifurcation with Old River	85.1	37.808	-121.327	4185838	647272
Mossdale Bridge	89.6	37.786	-121.307	4183410	649098
Sacramento River at Vernalis, station 11303500	111.8	37.676	-121.265	4171294	652968
Airport Way Bridge	111.9	37.676	-121.265	4171254	652984
Confluence with Stanislaus River	116.3	37.665	-121.242	4170130	655080
Maze Boulevard/Highway 132 Bridge	120.2	37.641	-121.229	4167489	656276
Old Fisherman's Club, boat launch	125.5	37.637	-121.193	4167049	659430
Finnegan Cut	127.5	37.626	-121.185	4165832	660166
Confluence with Tuolumne River	130.6	37.606	-121.174	4163701	661154
West Stanislaus Main Canal	131.1	37.605	-121.178	4163599	660851
Laird Slough	136.4	37.581	-121.165	4160866	662013
Grayson Road Bridge	139.9	37.563	-121.152	4158891	663185
Laird Park (upstream end)	142.0	37.560	-121.146	4158595	663752
Del Puerto Creek confluence	146.3	37.543	-121.116	4156782	666458
Las Palmas boat launch	154.4	37.498	-121.083	4151867	669490
East Las Palmas Avenue Bridge	154.9	37.494	-121.081	4151383	669669
Crow's Landing Road	169.0	37.431	-121.013	4144592	675768
Orestimba Creek confluence	171.8	37.422	-121.003	4143571	676686
Confluence with Merced River	187.6	37.349	-120.975	4135530	679341

Hydrology

Streamflow in the lower San Joaquin River is dependent on a combination of natural and engineered controls—climate, upstream dam regulation, diversions and return flows for agriculture and domestic water use, and reservoir releases for ecological purposes. The San Joaquin Basin has a drainage area of about 40,000 square kilometers (not including the Tulare Basin). Climate in the San Joaquin Valley is characterized by hot, dry summers and cool, mild winters, but annual precipitation is highly variable from year to year. Average annual precipitation in the valley floor is approximately 200 millimeters per year (mm/yr), but precipitation in the western slopes of the Sierra Nevada can exceed 1300 mm/yr (National Oceanic and Atmospheric Administration, 1985). The storms producing the greatest amount of precipitation result from extratropical cyclones (Dettinger and others, 2011), commonly referred to as “atmospheric rivers”.

Streamflow in all major tributary rivers is regulated by upstream dams, the largest of which are Friant Dam on the San Joaquin River, built in 1949; New Exchequer Dam on the Merced River, completed in 1967; New Don Pedro Dam on the Tuolumne River, completed in 1971; and New Melones Dam on the Stanislaus River, completed in 1978, which have capacities of 0.64 cubic kilometers (km³), 1.26 km³, 2.50 km³, and 2.96 km³, respectively. The latter three dams were built over smaller, existing dams—the Exchequer, Don Pedro, and Melones Dams—that had capacities of 0.35 km³, 0.37 km³, and 0.14 km³, respectively. The dams were built for water storage and flood control. They are able to capture much of the runoff from small- to mid-size storms, but do not have storage capacities great enough to capture all of the runoff from the largest storms.

Since the construction of dams in the San Joaquin River Basin, peak flows are usually less than natural flows before dam completion, as shown in the records of the peak annual streamflow for the four major tributary rivers (fig. 6). In figure 6, the completion date of major dam construction on each river is indicated, along with peak annual streamflow, recorded as discharge at selected streamgages, *upstream* from the dams. The upstream gages are mostly at the outlet of smaller watersheds; therefore, these records do not represent total natural flow or total inflow to the dam. Nevertheless, a comparison of the upstream and downstream peak annual streamflow records showed that many of the peak flows in the upstream hydrograph were absent or attenuated in the downstream hydrograph. Streamflows in Orestimba and Del Puerto Creeks are not regulated by dams.

The alterations in natural flow due to dam construction also were noticeable in daily mean discharge data from the USGS streamgage (San Joaquin River at Vernalis, Station 11303500; fig. 7). Daily mean streamflow records for water years¹ (WYs) 1924 and 1930–2014 were divided into two groups: pre-1967 and post-1967 (figs. 7A, B). The records were divided this way on the basis of the completion date (September 1966) of the New Exchequer Dam on the Merced River, which was the first of the “second-generation” dams constructed. In each plot, the records are grouped by hydrologic water-year type. The California Department of Water Resources (California Department of Water Resources, 2015) has categorized each WY from 1901 through 2014 for the San Joaquin Valley on the basis of unimpaired runoff volumes (State Water Resources Control Board, 1995). There are five types of water years in this hydrological classification system, ranging from wettest to driest: “wet,” “above normal,” “below normal,” “dry,” and “critical.” During a “wet” year, runoff is equal to or greater than 3.8 million acre-feet (MAF); for “above-normal,” runoff is greater than 3.1 MAF and less than 3.8 MAF; for “below normal,” runoff is equal to or less than 3.1 MAF and greater than 2.5 MAF; for “dry,” runoff is equal to or less than 2.5 MAF and greater than 2.1 MAF; and for “critical,” runoff is equal to or less than 2.1 MAF.

On average, the peak flows in post-1967 wet years were only slightly lower and arrived a little earlier than in pre-1967 wet years (the average of all wet years and average of all above normal years are shown as bold lines in figures 7A and 7B). During above-normal years, peak flows between April and June were notably lower after WY 1967, and during below-normal years, peak flows were essentially absent following dam construction, reflecting storage of water from snowmelt in reservoirs behind the dams during those water-year types. The altered hydrologic regime of the lower San Joaquin River following dam construction resulted in fewer and smaller high flows than before the dams were in place. This likely decreased the sediment transport capacity of the river as well.

During this study, data were collected over a range of water-year types—from wet in 2011 and dry in 2012 to critical in 2013 and 2014. The daily mean streamflow in the San Joaquin River, measured as discharge by the USGS streamgage at Vernalis (station 11303500), is shown for water years 2011–14 in figure 8. The minimum and maximum mean daily discharge (1923 to 2014) is also shown in figure 8 to compare the 4 water years of this study to the historical range.

¹A water year is the 12-month period from October 1 to September 30. It is designated by the calendar year in which it ends.

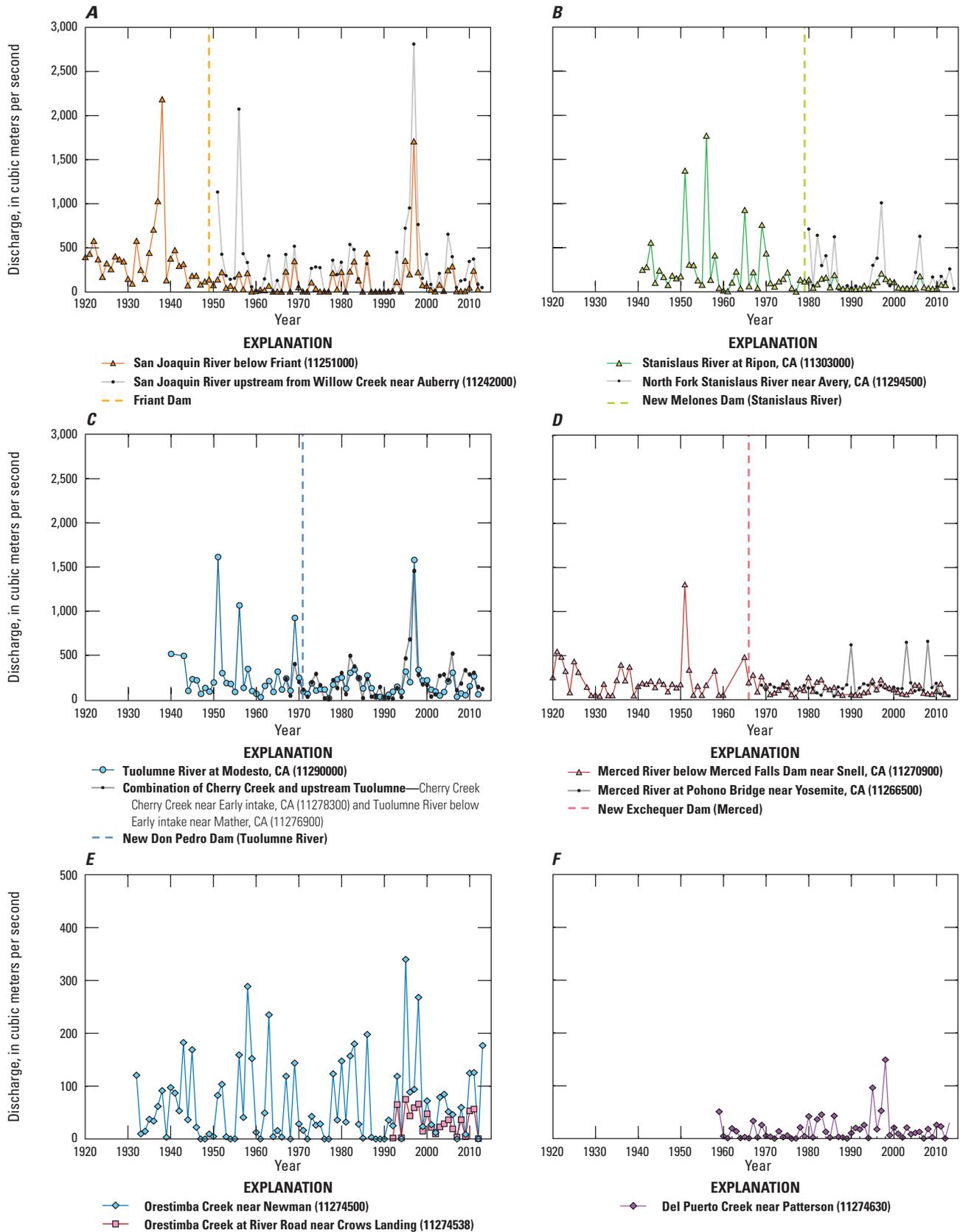


Figure 6. Annual peak streamflow, measured as instantaneous discharge at selected streamgages, and the time of dam completion, where applicable, for the following rivers and creeks in the San Joaquin Basin, California, 1920–2010: A, San Joaquin River; B, Stanislaus River; C, Tuolumne River; D, Merced River; E, Orestimba Creek; and F, Del Puerto Creek.

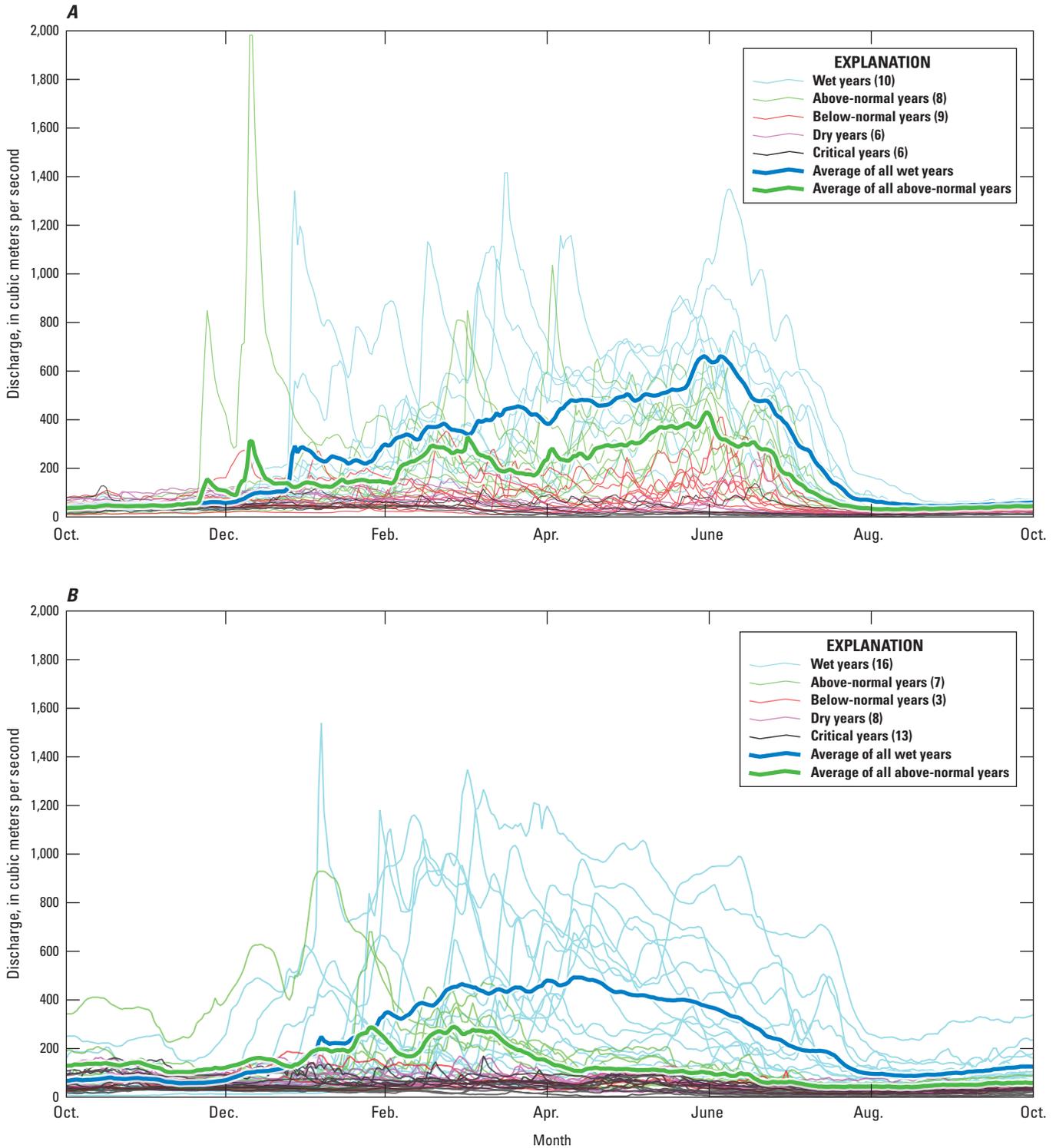


Figure 7. Daily average streamflow, measured as discharge at the U.S. Geological Survey streamgauge at Vernalis (station 11303500), for the lower San Joaquin River, California, categorized by water-year type (California Regional Water Quality Control Board, 1989) for water years *A*, 1924, 1930–67; and *B*, 1968–2014. Five types of water years were classified hydrologically on the basis of annual runoff (RO): in a “wet” water year, RO is equal to or greater than 3.8 million acre-feet (MAF); in an “above-normal” water year, RO is greater than 3.1 MAF and less than 3.8 MAF; in a “below normal” water year, RO is equal to or less than 3.1 MAF and greater than 2.5 MAF; in a “dry” water year, RO is equal to or less than 2.5 MAF and greater than 2.1 MAF; and in a “critical” water year, RO is equal to or less than 2.1 MAF. The number in parenthesis is the number of years included in that water year category.

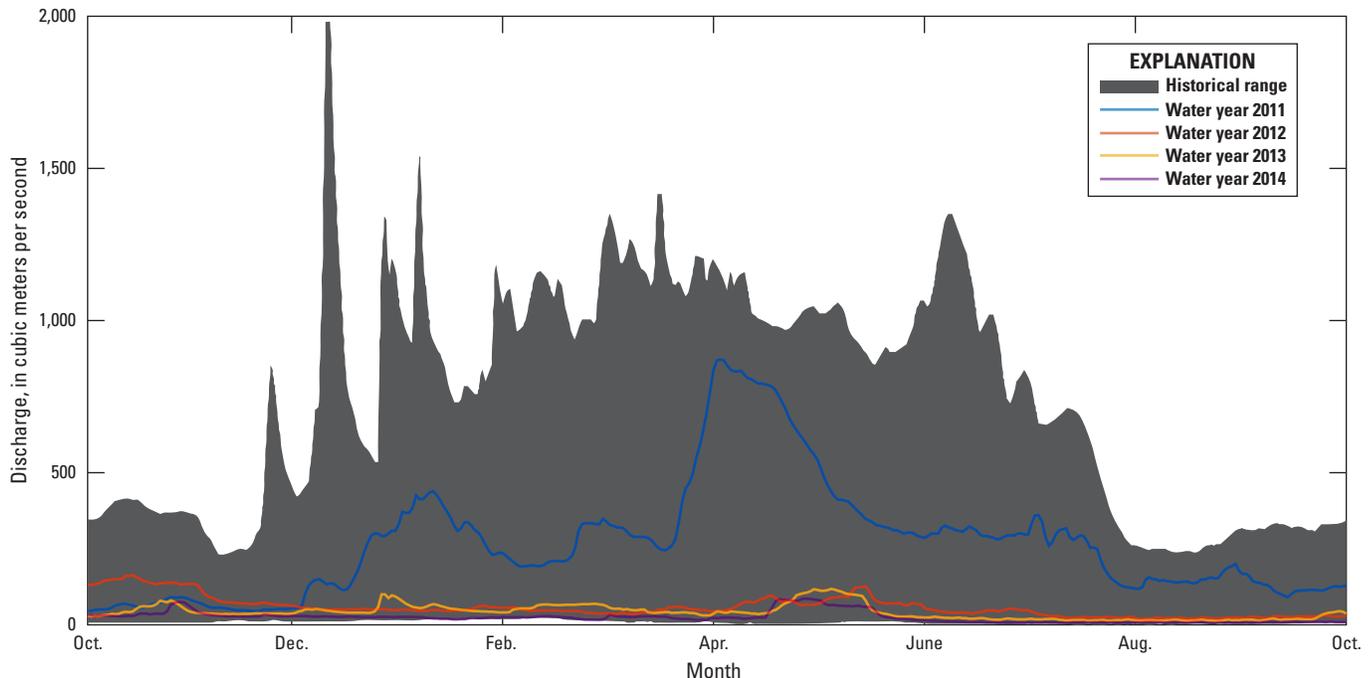


Figure 8. Daily mean streamflow, measured as discharge, during the study (water years 2011–14) relative to the historical range of discharge measured on the San Joaquin River, California, at the U.S. Geological Survey streamgage at Vernalis (station 11303500).

White Sturgeon Ecology

White sturgeon inhabit estuaries and rivers along the west coast of North America. They can live up to 100 years and often reach 2.5 meters (m) in length, although they can grow much larger (Moyle, 2002). White sturgeon are known to spawn in only three major river systems on the west coast: the Fraser, Columbia, and Sacramento–San Joaquin (Committee on the Status of Endangered Wildlife in Canada, 2012). They are generally considered diadromous because adults can live in fresh and salt water, but must spawn in freshwater. There is some evidence that adult white sturgeon stay in freshwater estuaries and do not venture out to the ocean (Veinott and others, 1999). There have also been cases, however, where individual sturgeon from the three major river systems were found in a different river system than the one in which they spawned (Nelson and others, 2004; Welch and others, 2006), and white sturgeon have been occasionally found in other systems, such as the Klamath River (Moyle, 2002). Adult Sacramento–San Joaquin white sturgeon spend most of their time in the San Francisco Bay estuary and migrate upstream in the spring to spawn (Schaffter, 1997). The juvenile fish can spend several years in the river channel before slowly migrating downstream to the estuary (Klimley and others, 2015).

Sturgeon are broadcast spawners that release eggs and sperm in fast moving water to disperse the eggs over a wide area (Conte and others, 1988). The location and timing of white sturgeon spawning are not well understood, but it is generally thought that the sturgeon respond to increases in

streamflow (Schaffter, 1997; Hildebrand and others, 1999; Paragamian and Wakkinen, 2011). They select areas in the channel that have the highest water velocity (Parsley and others, 1993; Parsley and Beckman, 1994; Billard and Lecointre, 2000). In the Kootenai River, for example, white sturgeon have only been found spawning in areas that have a mean water-column velocity of 0.8–2.8 meters per second (m/s; Parsley and others, 1993).

Sturgeon might choose areas with high water velocity because the bed material typical of such locations can increase the survival of eggs and larva. The type of bed material is important to white sturgeon because when the eggs make contact with freshwater, the eggs' outermost layer forms a thin adhesive coating that allows the eggs to adhere to the bed material (Conte and others, 1988). Eggs adhering to clean, coarse sediment (such as gravel or cobble) have better oxygenation (Kock and others, 2006). The eggs can adhere to other surfaces, such as bedrock, boulders, woody debris, submerged aquatic vegetation, or sand (Parsley and Kofoot, 2013). If the eggs are buried under as little as 2 millimeters (mm) of fine-grain sediment, however, mortality increases (Kock and others, 2006; Paragamian and others, 2009; Parsley and Kofoot, 2013). Viable eggs with a thin coating of sand have been documented in the lower San Joaquin River by Jackson and Van Eenennaam (2013) and in the Kootenai River by Paragamian (2012). Eggs can still develop if they land on sand-size bed material, but because sand is often highly mobile in a river, it is more likely for the egg to be buried and not survive.

The bed-material type can also play an important role in survival at the larval life stage of sturgeon. Larvae hide from predators in the interstitial spaces of gravels and cobbles, but tend to drift over surfaces such as sand or cobble embedded with sand (McAdam, 2011), which can lead to increased predation (Gadomski and Parsley, 2005). High water velocity (1.2–6.0 m/s) and channel depths greater than 1.8 m are thought to be characteristics of the preferred habitat for white sturgeon spawning (Gard, 1996).

Dams on river systems can adversely affect sturgeon populations (Rochard and others 1990; Parsley and Beckman, 1994). Dam construction can result in disconnection of habitat or isolation of individuals or subpopulations. Flow regulation from dams also cause downstream effects associated with diminished peak flows, such as seasonal decreases in water velocity, water depth, supply of coarse bed material, and sediment transport; thus, a long-term downstream effect of flow regulation by dams is a change in bed-material composition (Kondolf, 1997). Diminished peak flows could have affected the bed-material composition in many traditional white sturgeon spawning grounds. In the Kootenai River, for example, biologists have repeatedly observed that white sturgeon spawn in sand-bedded areas where cobble and gravel are buried beneath a meter or more of sand (Barton, 2004). Paragamian and others (2009) suggested that flow regulation at the upstream dam of the Kootenai River altered the sediment-transport regime and caused deposition of sand over traditional spawning areas, which sturgeon were selecting because of water velocity rather than bed-material type. Changes in land use and flow regulation due to upstream dam operations have notably altered the flow regime in the Sacramento–San Joaquin River system.

White Sturgeon Monitoring in the San Joaquin River by the U.S. Fish and Wildlife Service

In the spring of 2011, the USFWS used egg mats at eight locations to monitor spawning activity by white sturgeon (fig. 9B). At Rkm 142, 15 eggs were found on April 25, 2011, and another 4 eggs were found on April 28, 2011 (Gruber and others, 2012). The following year, egg mats were deployed at five locations: Rkm 115.2, 126.4, 137.6, 139.9, and 139.8 (fig. 9C). Viable eggs were found at all five locations between March and May 2012 (Jackson and Van Eenennaam, 2013). The river discharge on dates of the 2011–12 USFWS egg collection and of the USGS data collection are shown in figure 9A. Additional details about the USFWS white sturgeon monitoring are in Jackson and others (2016).

Study Methods

The equipment and methods used to collect data and samples to document the physical characteristics (water velocity, water depth, and bed-material sediment size) of the lower San Joaquin River are described in this section. Data generally were collected in the spring of each year from 2011 to 2014.

Vertical and Horizontal Datum

Elevation data were referenced to the North American Vertical Datum of 1988 (NAVD 88). Horizontal data are reported in grid coordinates of the Universal Transverse Mercator (UTM) projection, using the North American Datum of 1983 (NAD 83); the study area is in UTM zone 10 north.

River-Centerline Convention

Data and locations in this report are often referenced to a river-centerline coordinate system. A river centerline is a continuous line starting at the receiving water body, and it generally follows the middle of the channel, as delineated during typical low-flow periods. A river-centerline coordinate system measures distance along the centerline. Although some early USGS topographic maps show centerline coordinates in river miles, the course of the river has changed substantially; therefore, a new river centerline and coordinate system were defined for this study.

To create the river-centerline coordinate system, first, the wetted area of the river was digitized from the 2012 National Aerial Imagery Program (NAIP) orthorectified aerial imagery of the lower San Joaquin River. Next, a continuous line was drawn in the upstream direction tracing the course of the river in the center of the wetted area. In areas containing multiple channels or large islands, the centerline was drawn through the center of the widest channel. Although the study period includes samples and data collected during 2011–14, the channel configuration did not appear to change much in aerial imagery acquired between 2011 and 2014.

In most coordinate systems, distances are measured from the origin, or zero point. For the river centerline, stationing begins in the center of the channel of the downstream receiving water. For the San Joaquin River, the downstream receiving water is the Sacramento River; therefore, the junction of their centerlines was designated as the river kilometer origin, or Rkm 0. The geographical coordinates for each station at 1-Rkm intervals on the centerline of the lower San Joaquin River (Rkm 0–187) are provided in appendix table 1–1 in this report and are available online (Marineau and others, 2016). The river kilometer of selected landmarks and geographical features referenced in this report are provided in table 2.

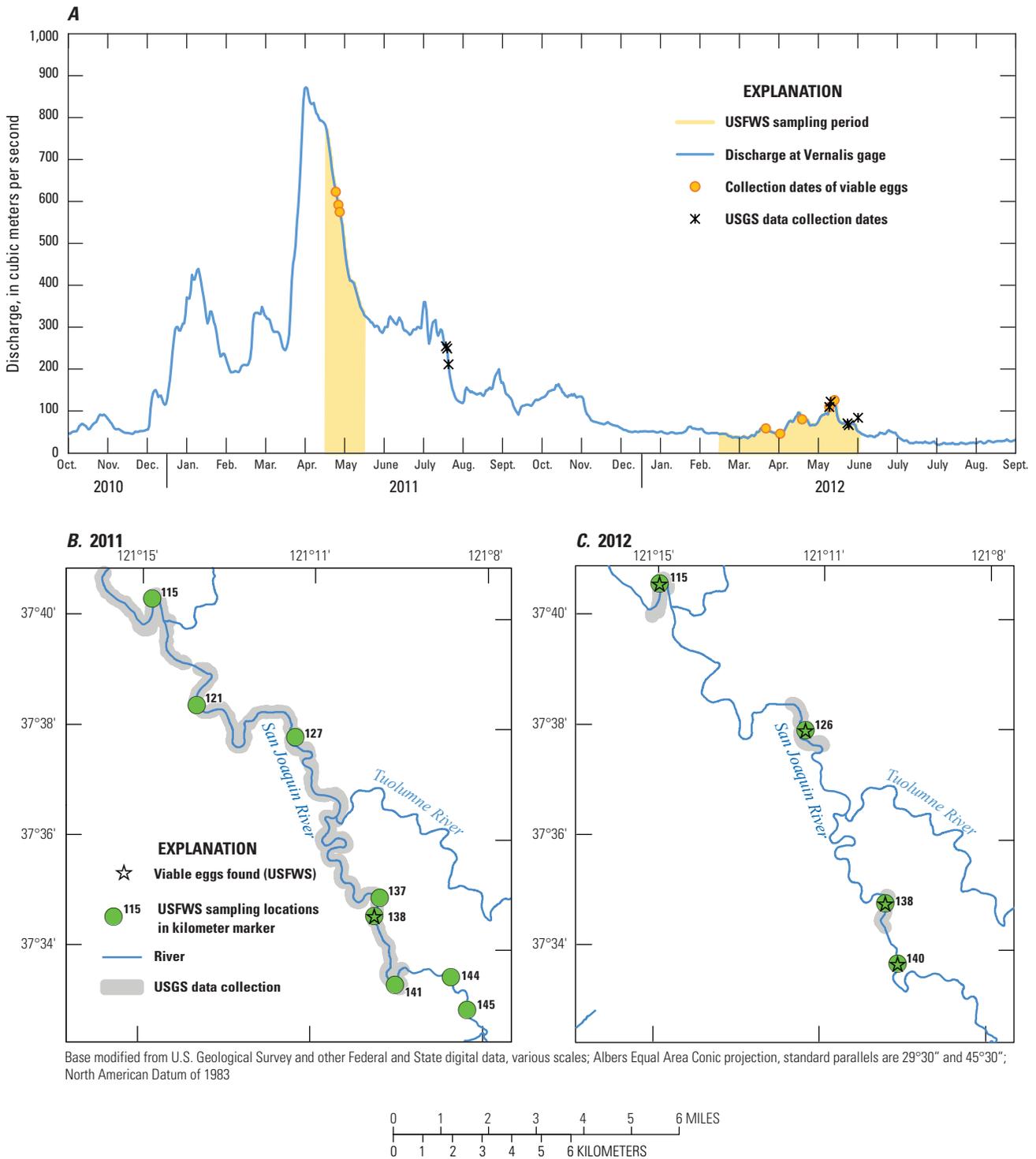


Figure 9. Dates and locations of the U.S. Fish and Wildlife Service (USFWS) sampling and U.S. Geological Survey (USGS) data collection, 2011–12: *A*, Hydrograph of discharge for San Joaquin River at USGS streamgauge at Vernalis, California (station 11303500) and discharge on the dates when viable white sturgeon eggs were collected by the USFWS and the USGS collected data. The USFWS sampling locations and extent of the USGS data collection; *B*, 2011; and *C*, 2012.

Equipment

The following is a list of equipment used during the data collection and sampling.

- 19-foot (ft) Koffler aluminum boat with outboard jet motor.
- 9.4-ft West Marine high-density polyethylene boat with 4-horsepower outboard motor.
- Odom™ multibeam echosounder system.
- SonTek® RiverSurveyor™ M9 acoustic Doppler current profiler (ADCP).
- SonTek real-time kinematic global positioning system (RTK-GPS).
- Teledyne RD Instruments Rio Grande™ ADCP.
- Field laptop with SonTek HydroSurveyor™ and WinRiver II™ software.
- Trimble R10 Global Navigation Satellite System (GNSS) receiver mounted to the boat used with a real-time network global positioning system (RTN-GPS).
- Trimble TSC3 hand-held controller, used to program the GNSS receiver and to record data.
- Humminbird™ 1198c SI Combo side-scanning sonar.
- US BMH-62 bed-material sampler.

Additional details (specifications, operation, and so forth) on the multibeam echosounder system, the ADCP equipment, and the side-scanning sonar follow.

Multibeam Echosounder System

High-resolution bathymetric data were collected in all four subreaches in 2011 using an Odom™ multibeam echosounder system. The multibeam system included a sonar head that operated at 240 kilohertz (kHz) and collected soundings over a 120-degree swath. The geolocation of the sonar head was determined with a global positioning system (GPS) receiver, which also recorded the heading or direction of travel. A motion reference unit (MRU) mounted directly behind the sonar head was used to precisely measure the orientation of the sonar head and the angle of the sonar beams. A sound-velocity sensor was mounted near the sonar head to measure the speed of sound at the head, data needed to form the received beam correctly. Data streams from the GPS, MRU, sound-velocity sensor, and sonar were integrated and processed using a computer running the hydrographic survey software. The system was mounted, tested, and calibrated in the field on July 25, 2011. The calibration involved determining the precise pitch, roll, and yaw angle offsets each time the system was mounted to a boat.

Acoustic Doppler Current Profiler

Water velocity and depth measurements were collected using two types of boat-mounted ADCPs. In 2011, 2012, and 2014, data were collected using a SonTek® M9 ADCP, and in 2013, data were collected using a Teledyne RD Instruments, Inc. (RDI), Rio Grande™ ADCP. Depth-averaged velocity was calculated from the vertical profiles measured by the ADCPs.

In general, both instruments collect the same type of data (water velocity, depth, and instrument position). The SonTek M9 is equipped with one center-mounted, downward-facing transducer (0.5 megahertz, or MHz), which measured the distance to the streambed, and two sets of four, equally spaced transducers mounted around the center transducer at an outward-slant angle. The two sets of transducers collect measurements at higher frequencies of 1.0 MHz and 3.0 MHz, respectively. The M9 ADCP automatically alternated between each set of transducer heads according to the detected water depth when in automatic adjustment mode. The RDI Rio Grande has a single set of four slant-mounted transducers operating at 1.2 MHz, with no vertical beam.

Horizontal positions were measured using a differential GPS, an RTK-GPS, or an RTN-GPS. In 2011, coordinates were measured using a differential GPS. In 2012, a SonTek RTK-GPS was used. As deployed, this RTK-GPS system only provided positions in a local reference system. In 2013, an RDI, Inc., differential GPS was used with the RDI Rio Grande™ ADCP. Finally, in 2014, a Trimble R10 RTN-GPS was used while taking measurements with the SonTek M9 ADCP.

Each ADCP used for this study was mounted to the side of the boat using an aluminum bracket, such that the instrument was submerged to a depth of 35 centimeters (cm) and was calibrated daily. Depth-averaged velocity was computed by RiverSurveyor Live for the SonTek M9 ADCP and WinRiver II for the RDI Rio Grande ADCP.

Side-Scanning Sonar

We used a Johnson Outdoors Marine Electronics, Inc., Humminbird™ 1198c SI Combo side-imaging sonar to collect side-scan sonar data. This instrument features two side-looking beams, one downward-looking beam, and a GPS receiver. The transducer was mounted to an adjustable-depth, hinged, aluminum bracket that was fixed to the transom of the motorized boat. The transducer was typically lowered to a depth of about 70 cm to avoid interference from “prop wash” from the primary boat engine; however, it was often raised up to a depth of only about 25 cm in shallow areas. At these times, the primary motor was raised, and only the smaller boat engine was used for propulsion. The side-imaging transducers operate at frequencies of 455 kHz or 800 kHz, whereas the downward-looking beam operates at 83 kHz or 200 kHz. The frequency of the side-imaging transducers was set to 455 kHz, as recommended by Kaeser and others (2012). Maximum range of the side-imaging transducers was adjusted in the field

on the basis of channel width to obtain optimal coverage and resolution of the entire channel. This distance was usually set at 20–30 meters.

Data Collection

This section describes the methods used to collect data and analyze samples for this study. Water depth was measured by soundings using a multibeam sonar and an ADCP. Bedform imagery was collected using a side-scan sonar. Efforts to map streambed material using the side-scan sonar images and ADCP backscatter-intensity data are discussed in this section. Bed-material sampling and processing are also described in this section.

Water Depth and Velocity Measurements using Acoustic Doppler Current Profiler, 2011–14

The locations of hydroacoustic data collection for each year are listed in [table 3](#) and shown in [figure 10](#). Locations of data collection varied from year to year, based

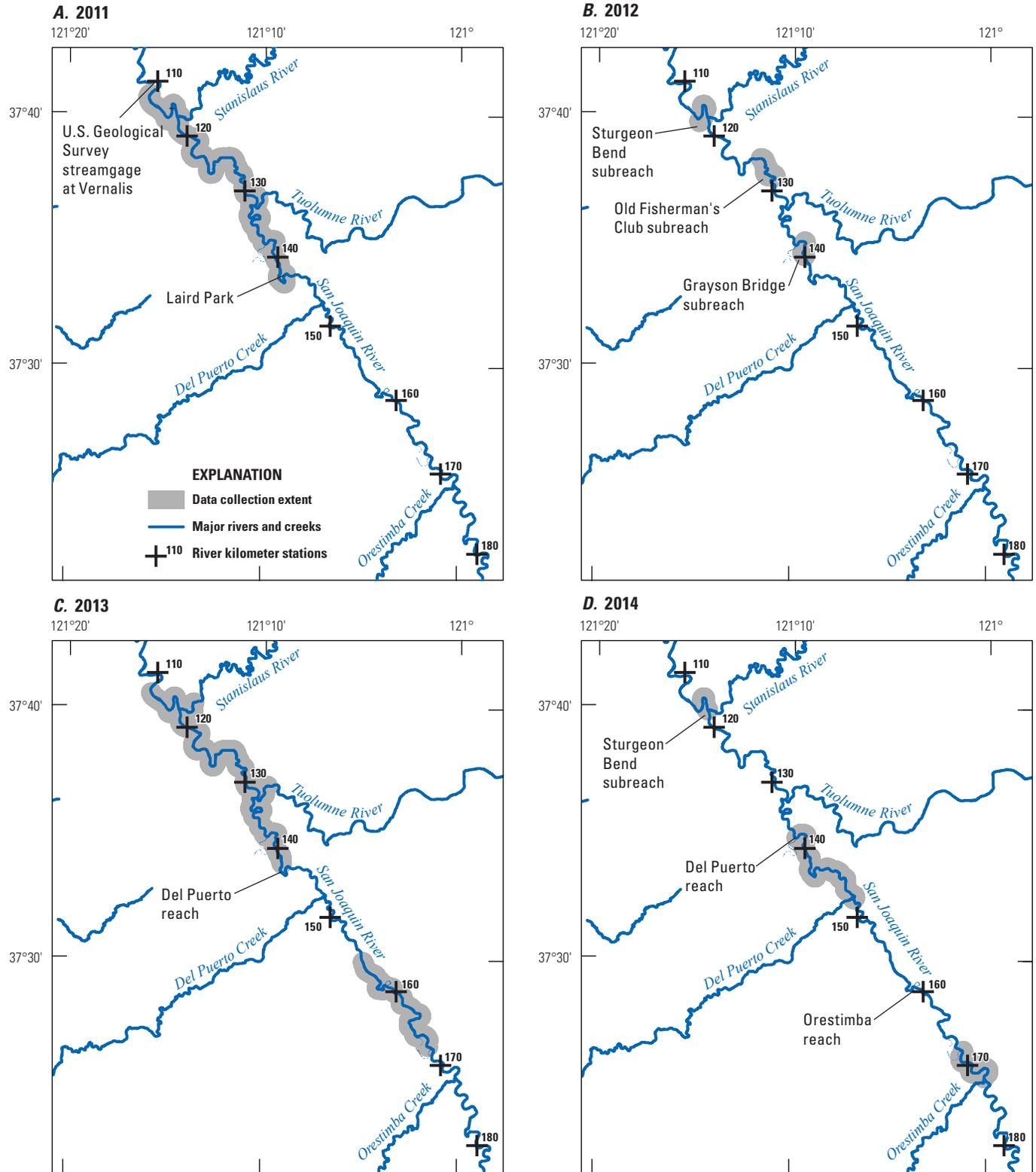
on conversations with USFWS biologists and as more information became available about the geomorphology of the river. In 2011, ADCP data were collected between Laird Park and the USGS streamgage at Vernalis (station 11303500) at Rkm 111.8 ([fig. 10A](#)). In 2012, data collection focused on the three subreaches: Sturgeon Bend, Old Fisherman’s Club, and Grayson Bridge ([fig. 10B](#)). In 2013, although low-flow conditions limited access to many areas, the data were collected for the same reach as in 2011, for small portions of the Tuolumne and Stanislaus Rivers, and for part of the Orestimba reach ([fig. 10C](#)). In 2014, a smaller boat was used to improve access, but low-flow conditions still presented a challenge and limited access to many parts of the river. Data collection in 2014 included the Sturgeon Bend subreach, the Del Puerto reach, and a small part of the Orestimba reach ([fig. 10D](#)).

Depth and velocity data were collected using two different approaches: channel cross sections and longitudinal profiles. A combination of channel cross sections and longitudinal profiles were collected in 2011 and 2012, and also in limited areas in 2014. In 2013, only longitudinal profiles were collected in order to cover large areas quickly.

Table 3. Summary of hydroacoustic data collection sites, including daily mean streamflow, measured as discharge, in the reach during data collection and the exceedance probability of that discharge.

[Hydroacoustic data collection sites (and Rkm locations) refer to the San Joaquin River, California, except where otherwise indicated. **Abbreviations:** mm/dd/yyyy, day/month/year; m³/s, cubic meters per second]

Sites and river kilometers (Rkm) of data collection	Date (mm/dd/yyyy)	Discharge (m ³ /s)	Exceedance probability (percent)
2011			
Stanislaus reach (Rkm 111.5–116.3)	07/19/2011	254	13
Tuolumne reach (Rkm 116.3–118.2)	07/19/2011	197	10
Tuolumne reach (Rkm 118.7–130.5)	07/20/2011	174	11
Del Puerto Creek reach (Rkm 130.5–141.8)	07/21/2011	58	15
2012			
Grayson Bridge subreach (Rkm 137.3–139.9)	05/10/2012	38	24
Old Fisherman’s Club subreach (Rkm 125.2–127.5)	05/24/2012	30	36
Sturgeon Bend subreach (Rkm 114.0–115.7)	05/25/2012	66	44
2013			
Sturgeon Bend subreach (Rkm 111.9–119.8)	04/22/2013	76	38
Stanislaus River (Rkm 0 to 1.9)	04/22/2013	42	18
Grayson Bridge subreach (Rkm 130.3–139.8)	04/24/2013	17	64
Old Fisherman’s Club subreach (Rkm 116.2–130.9)	04/25/2013	31	28
Tuolumne River (Rkm 0–1.8)	04/25/2013	19	43
Orestimba reach (Rkm 154.4–166.7)	04/26/2013	15	70
2014			
Del Puerto Creek reach (Rkm 136.4–146.6)	05/12/2014	5	95
Orestimba reach (Rkm 167.2–171.8)	05/13/2014	5	95
Sturgeon Bend subreach (Rkm 114.9–116.3)	08/22/2014	6	99



Base modified from U.S. Geological Survey and other Federal and State digital data, various scales; Albers Equal Area Conic projection, standard parallels are 29°30" and 45°30"; North American Datum of 1983

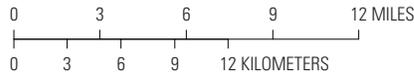


Figure 10. Locations where water depth and velocity data were collected, by year, in the lower San Joaquin River, California, 2011–14: A, 2011; B, 2012; C, 2013; and D, 2014.

Bathymetry Mapping using Multibeam Sonar, 2011

The multibeam sonar survey was completed in 3 days, July 26–28, 2011, along the reach of the lower San Joaquin River between Laird Park and the USGS streamgaging station 11303500, San Joaquin River near Vernalis. Detailed bathymetric maps, based on 2- by 2-m grid spacing, were created for the reach using HYPACK® hydrographic survey software. This resolution was selected to guide efficient collection of the bathymetric points and to produce maps sufficiently detailed to capture the variable topography of the river bed. The only year that had flows sufficiently high to justify using multibeam sonar was 2011; flows in subsequent years were too low for a swath-mapping technique like multibeam sonar to be more efficient than ADCP mapping.

Bed-Material Mapping using Acoustic Backscatter in 2012 and Side-Scan Sonar in 2013

In 2012, the ADCP backscatter data were analyzed using the methods described by Shields (2010) and compared to bed-material samples. An ADCP operates by transmitting high-frequency acoustic pulses in four directions, and then it records the time and intensity of the reflection of those pulses. The instrument uses the travel time and the frequency shift to compute water velocity, but it also stores details of the amplitude of the return energy (backscatter intensity). Harder, denser bed material, such as gravels, cobbles, bedrock, and consolidated clay (hardpan), can reflect more of the incident acoustic energy than do softer sediments, such as unconsolidated clay, and porous sediments, such as sand, that absorb the energy. The correlation between ADCP acoustic backscatter intensity at the bed and measured particle size was poor (possibly reflecting the coarse resolution of the backscatter intensity from vertical averaging over 0.5–1.0-m cells), however, and was not explored further.

In 2013, two methods were evaluated to identify sediment type using side-scan sonar data. A side-scan sonar is similar in some aspects to the ADCP, except instead of measuring velocity, it uses a wide-angle beam to record a swath of high-resolution backscatter intensity. The high-resolution backscatter-intensity data are stored as an image slice, where the digital number of the image pixel represents the relative intensity of the backscatter. These image slices can be post-processed to correct for boat position and heading and then mosaicked to form a map.

The first method used to analyze the side-imaging data was to mosaic the images and view them manually in an attempt to identify bed composition (Kaesler and others,

2012). Two commercially available software programs were compared to evaluate their ability to post-process the raw data files: Dr. Depth (Johnson Outdoors Marine Electronics, Inc.) and Sonar TRX-SI-Pro (Leraand Engineering, Inc.). Both programs were able to mosaic the data, to correct for boat position and heading, and to provide output images of the downward-looking beam and side-looking beams (figs. 11, 12).

The second method tested was to process the data from the downward-looking beam to determine hardness on the basis of the intensity of the acoustic backscatter. The Dr. Depth program was capable of providing an output of “hardness,” although no explanation of the calculation for hardness was provided in the documentation. This method was used previously by others to map the bed material in the James River, Virginia (Austin, 2012). In theory, areas with high backscatter intensity contain gravel, cobble, or bedrock, whereas areas with low backscatter contain sand, silt, and clay.

Areas that were identified as likely to have sand, silt, or clay on the basis of side-scan imagery and also had low backscatter intensity were thought to have sand or finer bed material. This fine bed material made up the majority of the survey areas and also reflected that reported by previous literature (Hall, 1886; Mussetter Engineering, Inc., and others, 2000). Areas that had medium or high backscatter intensity, or where the bed-material type could not be easily identified from the side-scan imagery, were thought to potentially contain coarse bed material, such as gravel or cobble. There were not many of these areas, so the bed material at all such sites was sampled in the field. The areas with high backscatter intensity were associated with gravel; cobble rip-rap; hardpan clay; or, occasionally, sand in shallow (less than 0.5 m) water. Because the San Joaquin River is predominately a sand-bedded river and a primary interest of this study was to identify the few areas with gravel or cobble bed material, the combination of the two side-scan methods worked well to target sampling in locations with coarser bed materials.

Tributary Longitudinal Elevation and Slope Profiles

Channel gradient (or slope) in a river can provide some indication of the forces, such as shear stress at the riverbed, available to transport sediment relative to the size of that sediment. River reaches that have steep slopes (such as in mountains) tend to have higher bed shear stresses and transport a wide range of sediment sizes (boulders to silt), whereas river reaches with low slopes (such as on a valley floor) tend to have lower stresses and can only transport finer sediment (such as sand and silt).

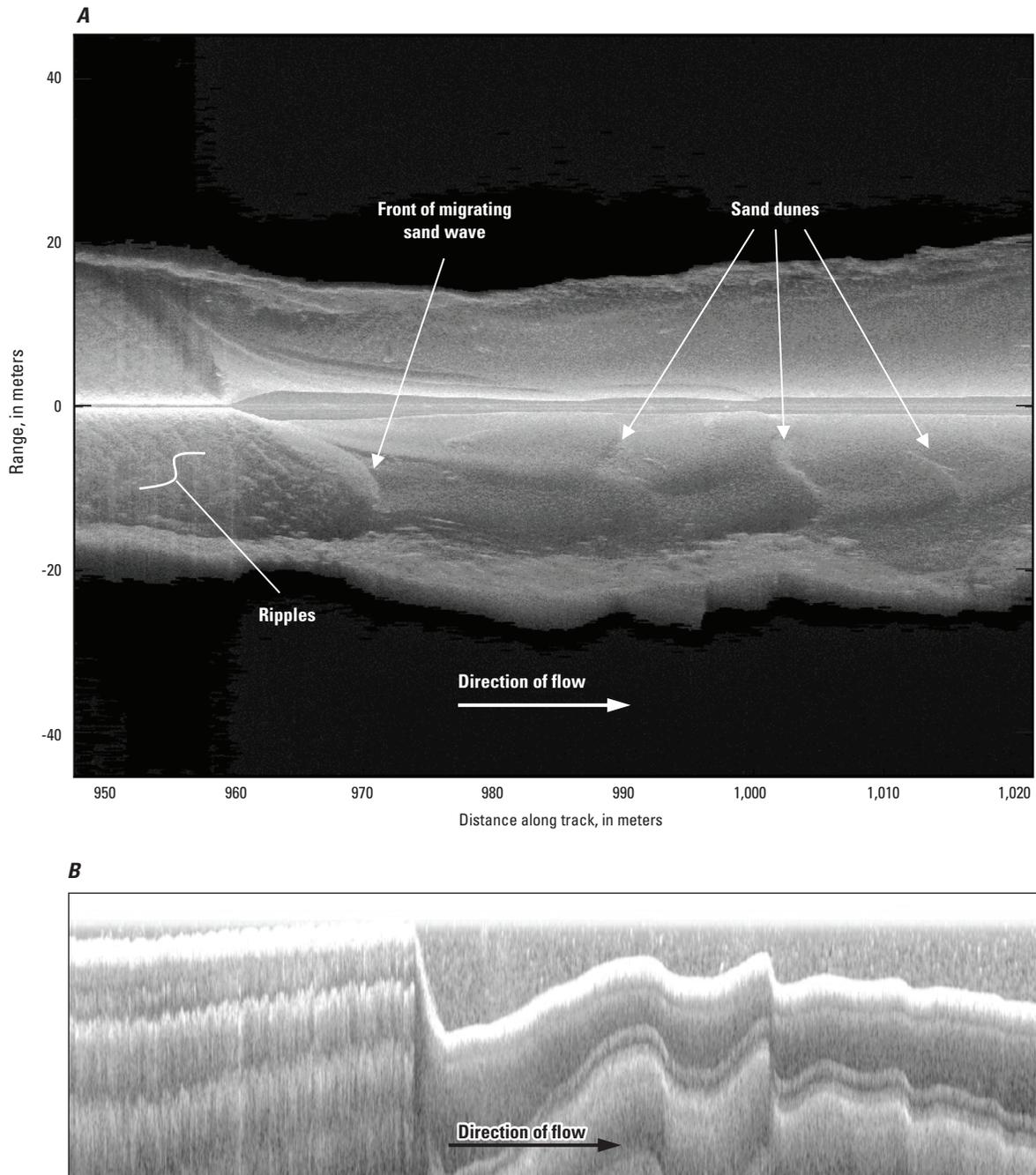


Figure 11. Examples of side-scan sonar from a sand-bedded section in the Grayson Bridge subreach of the lower San Joaquin River, California, April 23, 2013: *A*, plan view; and *B*, profile view showing small ripples translating up the stoss slope on an advancing sediment wave (left) and large migrating sand dunes (right).

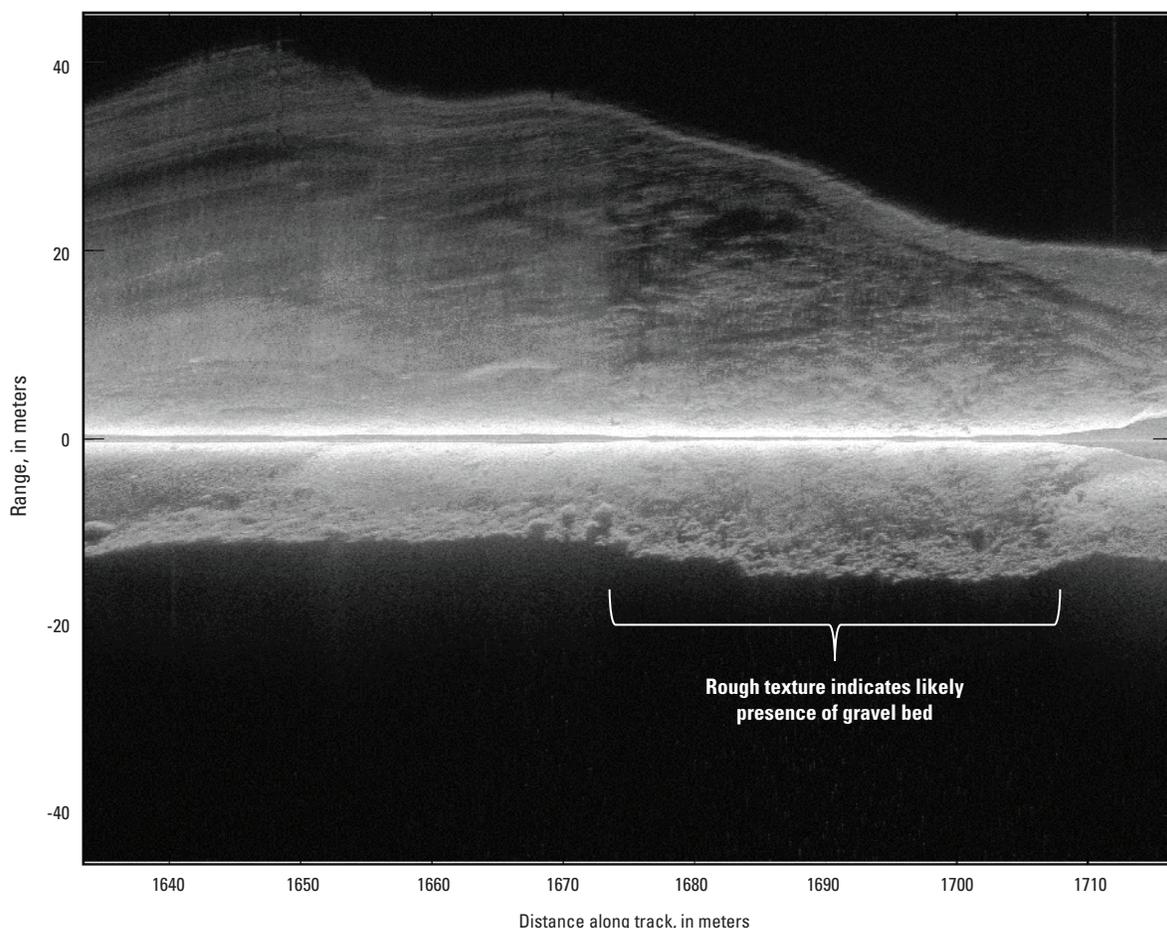


Figure 12. Plan view of side-scan sonar of a gravel-bedded section in the Grayson Bridge subreach of the lower San Joaquin River, California, April 23, 2013. The rough texture of the bed indicates likely presence of gravel, which was confirmed by collection of bed-material samples in this area.

We created longitudinal elevation profiles of the San Joaquin River, the three major east-side tributary rivers (Stanislaus, Tuolumne, and Merced), and two west-side tributary creeks (Del Puerto and Orestimba) using elevation data from the National Elevation Dataset (U.S. Geological Survey, 2015a), which had a 30-m grid resolution, and centerlines from the National Hydrography Dataset (U.S. Geological Survey, 2015b). Although water-surface elevation was surveyed in some areas using the GPS-RTN, this dataset was not sufficient to generate maps for the entire study area in the lower San Joaquin River or for the tributary rivers. For each river or creek, we divided the river centerline into stations spaced at 100-m intervals. The elevation at each point was assigned by sampling the NED-derived elevation model using spatial analysis tools in ArcMap; each centerline station also was indexed to the distance upstream from the tributary mouth. We then derived channel gradient using the change in elevation between each 100-m station and smoothed the channel-gradient profile using a 10-km averaging window. The 10-km averaging window helps smooth over uncertainties

in elevation that result from the digital elevation model of the centerline not always reflecting the true, within-bank, water-surface elevation, but instead providing a sense of the general gradient of the river. This provided an appropriate dataset for comparing reach-averaged gradients among the reaches.

Bed-Material Sampling, 2012–14

Bed material was sampled using a US BM-54 and US BMH-60, which are cable- and handline-suspended versions, respectively, of a rotary-scoop bed-material sampler, following standard USGS protocols for use of these samplers (Subcommittee on Sedimentation, 1958; Edwards and Glysson, 1999; Davis, 2005) from a motorized boat. These types of samplers are designed to collect material from the top 5 cm of the bed. Samples were collected during 3 years of the study (2012–14), but were not collected during the first year of the study (2011). Approximate locations of bed-material samples collected in each water year are shown in [figure 13](#).

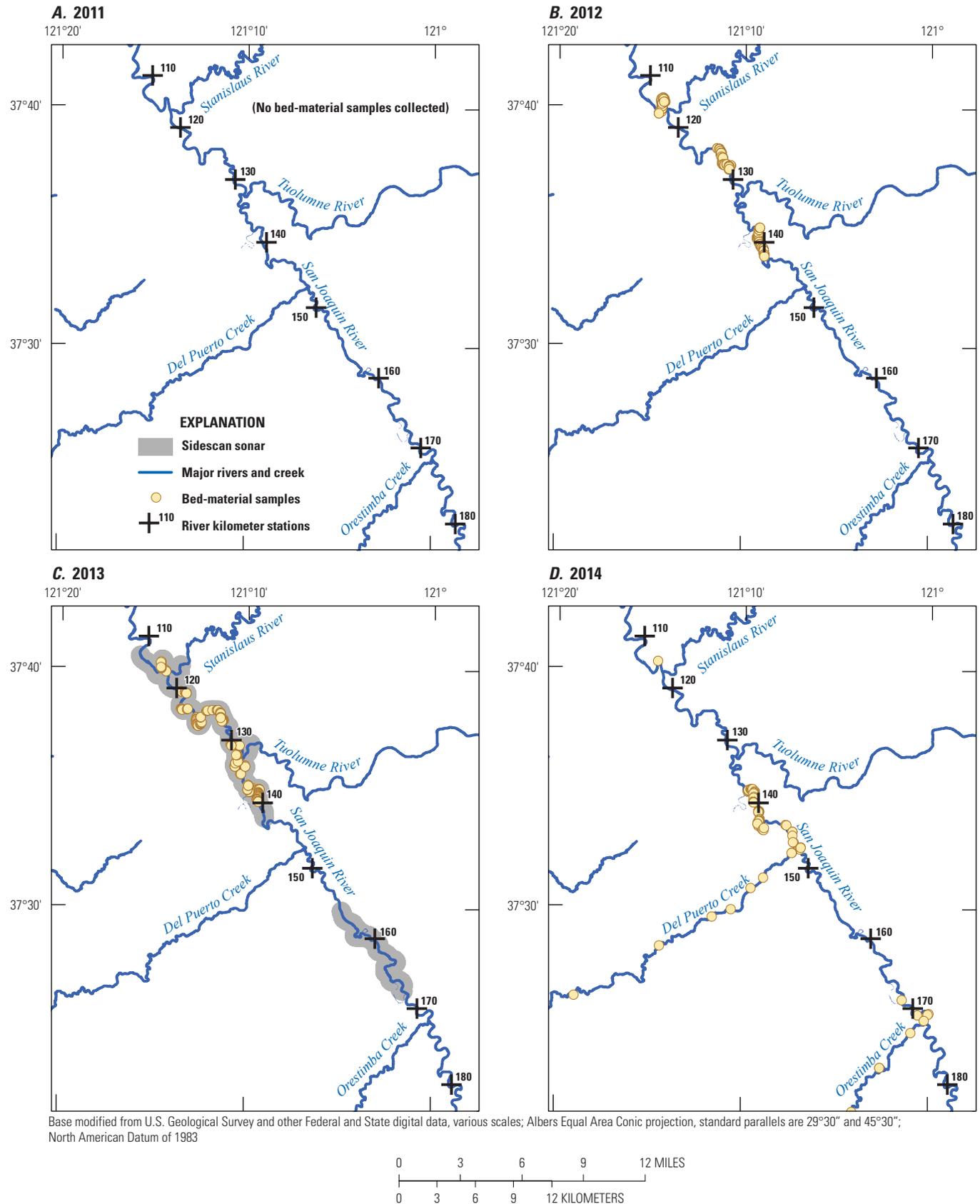


Figure 13. Sites where bed-material samples were collected in the lower San Joaquin River, Del Puerto Creek, and Orestimba Creek, California: *A*, 2011; *B*, 2012; *C*, 2013 (including the linear extent of 2013 side-scan sonar data collection); and *D*, 2014.

The locations of sample collection were recorded using the boat's on-board GPS navigation system, a mobile phone with a GPS software application, or by cross-referencing sample time with the GPS coordinates stored in the ADCP. The accuracy of the boat's on-board GPS was about plus or minus 2.5 m (Johnson Outdoors Marine Electronics Inc., 2014), and the mobile phone GPS was about plus or minus 8 m (Zandbergen, 2009). Additional information about sampling (location in channel, date and time, and depth of sample) was also recorded routinely.

In 2012, transects were established at the three subreaches to collect bed-material samples and ADCP measurements. At each transect, the channel was divided into five equal-width sections, and a bed-material sample was collected at the center of each section.

In 2013, the sampling protocol was modified to increase spatial coverage. First, longitudinal profiles of the streambed were collected as side-scan sonar images. These data were processed and reviewed to identify "areas of interest" using backscatter-intensity data and images (the methods are discussed in the previous section). "*Areas of interest*," for the purpose of this study, were areas that potentially had coarse bed material, as indicated either by high acoustic backscatter intensity or the side-scan imagery showing rough texture or heterogeneous surfaces (for example, [fig. 12](#)) that could potentially signal the presence of coarse bed material. At areas of interest, one or more bed-material samples were collected in the center of the channel. If an area of mostly coarse bed material was found, additional samples were collected upstream and downstream to identify the longitudinal extent of the coarse bed material. Based on the results from the 2012 bed-material transect samples, the sediment type did not vary much across the streambed laterally. Therefore, a single bed-material sample collected from the center of the channel was deemed sufficient to represent the bed material at that longitudinal station. On occasion, a sampler came up empty or only contained a small amount of clay material. This was assumed to indicate the sampler hit hardpan clay that the scoop could not penetrate.

In addition to the bed-material samples collected in the main channel, samples were collected from streambanks, gravel bars, and the west-side tributary creeks in 2014 to try to determine the likely source of the coarse bed material (gravel and cobble) in the lower San Joaquin River. Low-streamflow conditions in 2014 limited the ability to collect bed-material samples using traditional sampling techniques. In shallow areas (less than 1.2 m), a metal pole was used to prod

the streambed to determine if the bed was composed of soft material (sand and silt) or hard material (gravels and cobbles). If gravel or cobble was detected using the prodding method in a wadable area, then a grab sample was collected; if the area was too deep, the location was recorded, but no sample was collected.

Bed-Material Sample Preparation and Particle-Size Analysis

Bed-material samples were dried in an oven at 105 degrees Celsius. After samples were dried, clumps of sediment were disaggregated using a mortar and pestle. Particle-size distribution of the sample material was then determined using a combination of sieving and direct measurement.

Sieving of sediment particles smaller than 2 mm followed standard sieving methods (Guy, 1969). For the sieved portions, the fraction of particles finer than 0.063 mm, 0.125 mm, 0.5 mm, 1 mm, and 2 mm were weighed and recorded. If a sample contained particles larger than 2 mm, those particles were sorted by size manually using an SAH-97 hand-held particle-size analyzer (Potyondy and Bunte, 2002). The total weight of each size class was recorded and then used with the sieve data to determine the particle-size distribution.

The median, tenth-percentile, and ninetieth-percentile grain sizes were calculated using linear interpolation between logarithmically transformed particle-size classes. The percentages of fine sediment, sand, gravel, and cobble in each sample were also calculated. In this report, fine sediment is smaller than 0.063 mm; sand is from 0.063 to less than 2 mm; gravel is from 2 to less than 64 mm; and cobble is from 64 to less than 256 mm ([table 4](#)).

Table 4. Description and size range of sediment particles in particle-size categories used in this report.

[mm, millimeter; <, less than; >, greater than]

Particle-size category	Size range (mm)
Fine (silt/clay)	<0.063
Sand	0.063–2
Gravel	>2–64
Cobble	>64–256

Summary of Channel Characteristics, Stream Slope, and Bed Aggradation in the Lower San Joaquin River

The following section provides a summary of the results from this study related to (1) water depth and velocity, (2) bed material, (3) stream slope, and (4) aggradation. Water depth and velocity measurements provided information about general conditions in the lower San Joaquin River during these years, particularly the identification of deep scour holes and high-velocity areas. The lower San Joaquin River was nearly all sand-bedded, except for a series of gravel-bedded areas downstream from the Del Puerto Creek confluence and a small patch downstream from the Orestimba Creek confluence. The stream slope of the lower San Joaquin River and most of its tributaries is low, except in Del Puerto and Orestimba Creeks. Bed aggradation was observed in areas through repeated measurements, such as in the Sturgeon Bend subreach.

Depth and Velocity

Ranges of the approximate channel depth and depth-averaged velocity for the entire study area between 2011 and 2014 are shown in [figures 14](#) and [15](#). Most of the depth data presented here were collected using an ADCP; however, some depth data were collected using an Odem multibeam in 2011. The depth data from the multibeam and ADCP were combined for this section of the report. Detailed depth and velocity data are available online (Marineau and others, 2017).

Water Year 2011: Wet

The California Department of Water Resources (2015) classified 2011 as a wet water year. In the spring of 2011, a combination of heavy rainfall and rapid snowmelt caused streamflows in the lower San Joaquin River to increase over a few months, peaking at 875 cubic meters per second (m^3/s) on April 2, 2011 (recorded at *San Joaquin River near Vernalis, CA, Station 11303500*). Field data were collected July 19–21, 2011, when the streamflow was much lower (but still relatively high for the river), ranging from 196 to 256 m^3/s .

Because the highest flows during the study were in 2011, water depth and velocity in that year were higher than during all subsequent 3 years of data collection. The maximum depth recorded was 14.0 m, at Sturgeon Bend. Several deep pools were found in the river that year. In the 30-km reach (Rkm 111.5–141.8) surveyed in 2011, 21 pools deeper than 6 m were recorded. Of those, five were deeper than 9 m (not including Sturgeon Bend). Most of these deep areas were associated with sharp river bends, which can have complex and dynamic flow patterns, and were typically about one

channel width long. On average, there was one deep pool (greater than 6 m) for every 1.4 km of channel surveyed.

During the 2011 data-collection period, depth-averaged velocity exceeded 0.5 m/s in all river reaches measured, was very often more than 1.0 m/s, and exceeded 2.0 m/s in some areas. A visual inspection was made of the depth relative to velocity by plotting points where depth and velocity data were both available. Based on that plot, there was no apparent correlation between channel depth and depth-averaged velocity.

Water Years 2012–14: Dry and Critical

Water years 2012 and 2013–14 were dry and critical respectively, based on runoff in the San Joaquin Basin (California Department of Water Resources, 2015). During these years, the peak flow measured at the downstream end of the study area was about 80–90 percent less than it was in 2011 ([fig. 8](#)) and was 165, 119, and 87 m^3/s for 2012, 2013, and 2014, respectively (recorded at *USGS streamgage 11303500, San Joaquin River near Vernalis, CA*).

In water years 2012–14, depth and velocity were generally less than that measured in 2011. For instance, water depth measured at Sturgeon Bend decreased from 14.0 m in 2011 to 9.4 m in 2012, 5.8 m in 2013, and 3.8 m in 2014—a total decrease of about 10 m in 4 years. This decrease in depth was only partially related to the low water stage during those dry years (relative to bed features that have persisted from year to year); it primarily was caused by 8 m of sediment aggradation during the same period (discussed later in this report).

Depth-averaged streamflow velocity was typically less than 1.0 m/s during the dry and critical water years ([fig. 15](#)). There were isolated areas where depth-averaged velocity exceeded 1.0 m/s, particularly downstream from the confluence with the Stanislaus River. The higher velocity in the main stem downstream from this confluence appeared to be primarily due to reservoir releases at New Melones Dam on the Stanislaus River. Streamflow in the lower San Joaquin River upstream from the Stanislaus River confluence was about half of that recorded at the Vernalis gage.

Bed-Material Composition and Sources

Results of the bed-material size analysis are shown on remotely sensed imagery along with water depth or velocity measurements in [figures 16–21](#). Additional sediment samples were collected from dry stream beds or gravel bars ([figs. 22–27](#)). A plot of the 50th (median), 10th, and 90th percentiles of particle size in each sample by river kilometer along the lower San Joaquin River and up Del Puerto and Orestimba Creeks is shown in [figure 28](#).

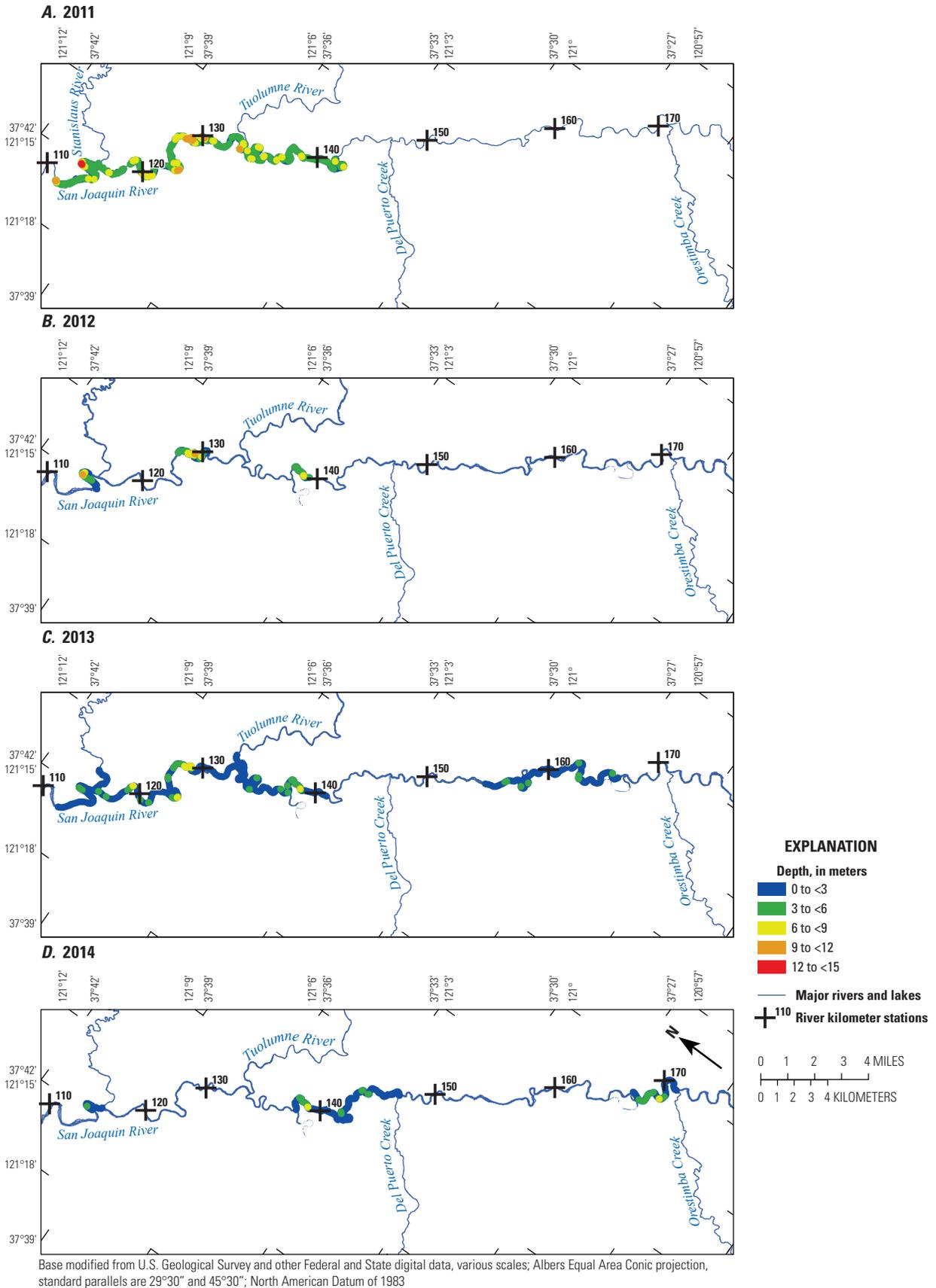


Figure 14. Water depth in the lower San Joaquin River, California: *A*, 2011; *B*, 2012; *C*, 2013; and *D*, 2014. Symbols for greater depth values are overlain on symbols for lesser values.

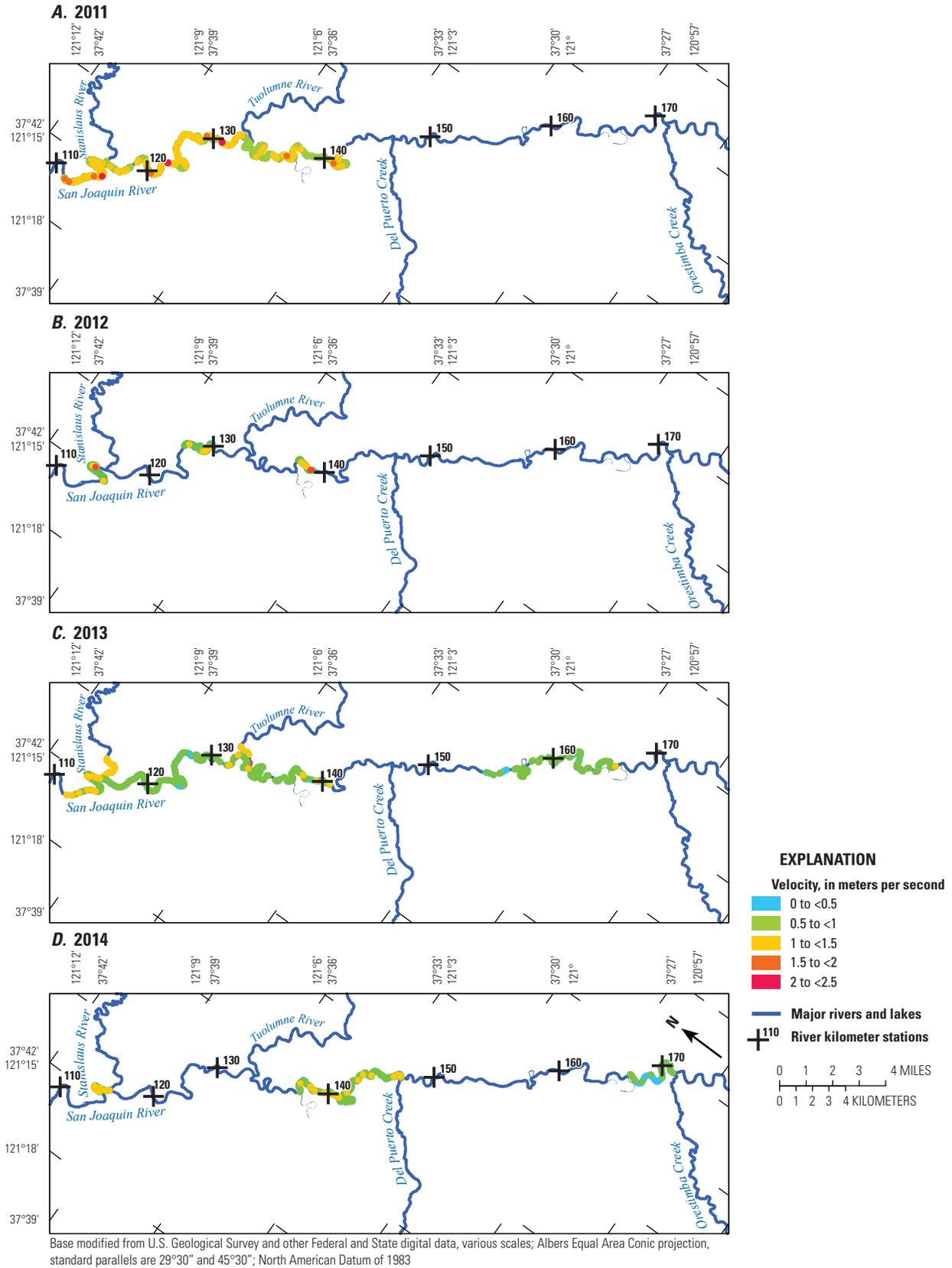


Figure 15. Depth-averaged velocity in the lower San Joaquin River, California: A, 2011; B, 2012; C, 2013; and D, 2014. Symbols for greater depth-averaged velocity values are overlain on symbols for lesser values.

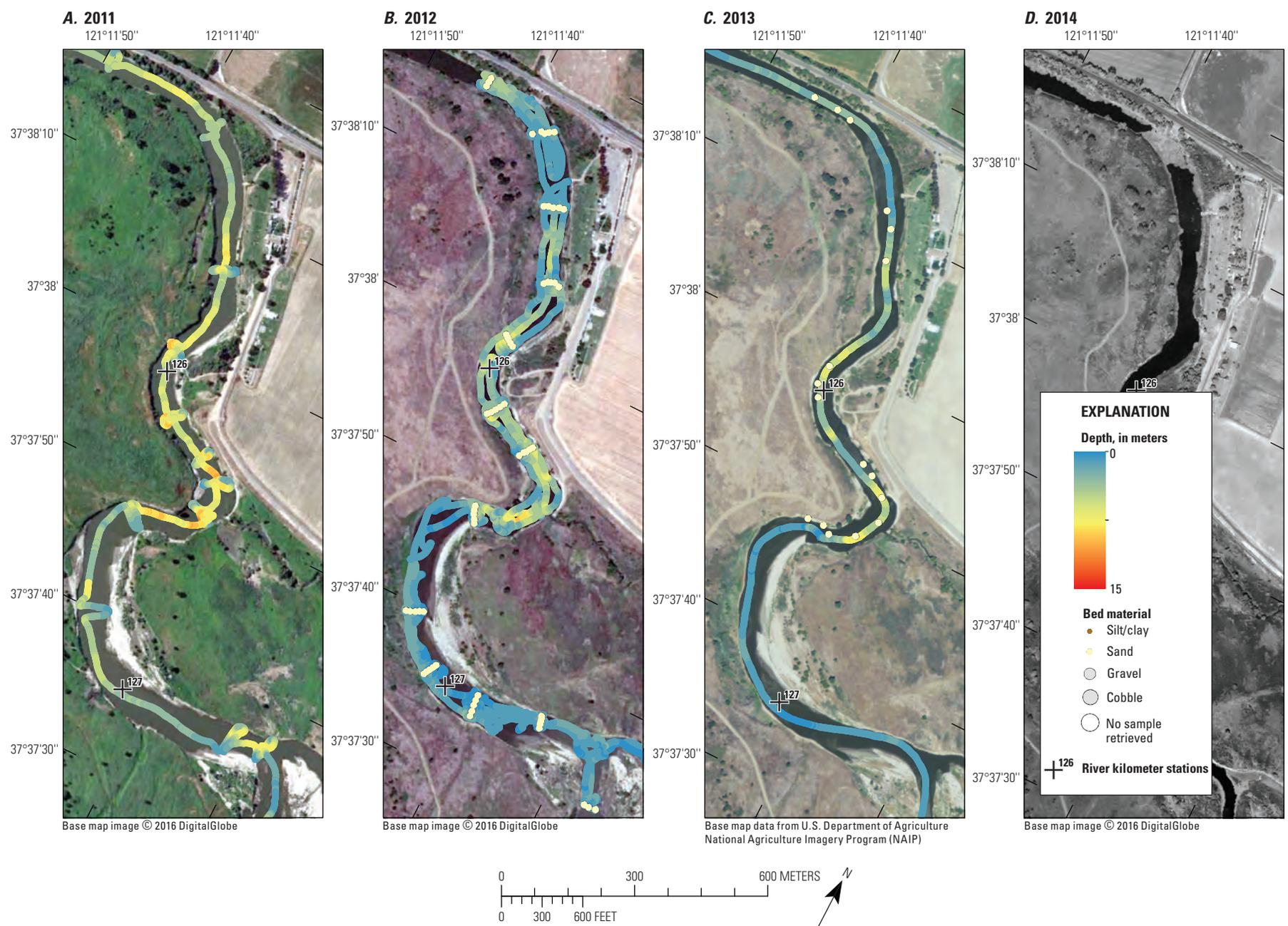


Figure 17. Water depth and bed-material size class in the Old Fisherman’s Club subreach of the lower San Joaquin River, California (river kilometers 125.2–127.5): *A*, 2011; *B*, 2012; *C*, 2013; and *D*, 2014. No depth data or bed-material samples were collected in 2014 from this subreach.

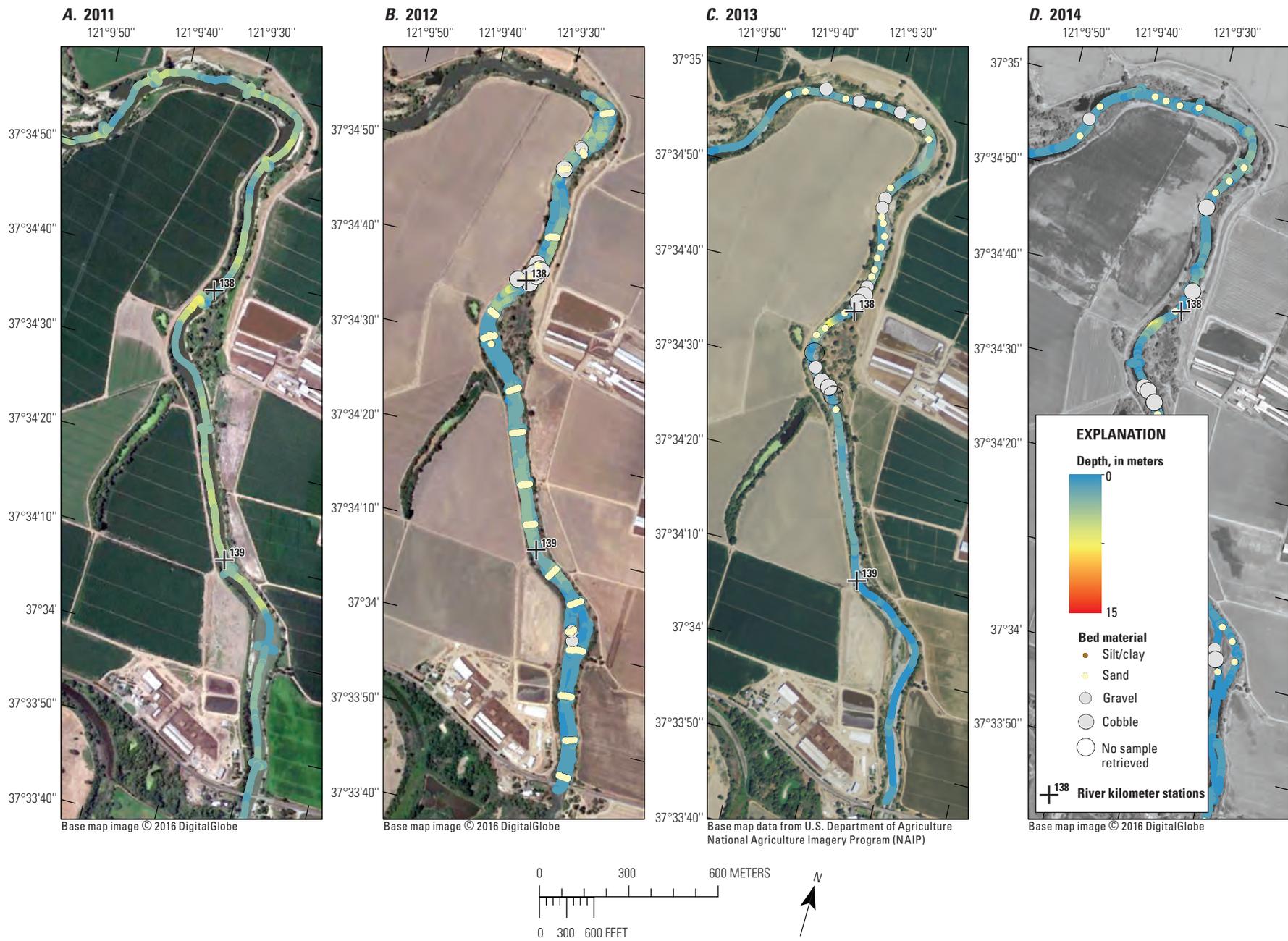


Figure 18. Water depth and bed-material size class in the Grayson Bridge subreach of the lower San Joaquin River, California (river kilometers 137.3–139.9): A, 2011; B, 2012; C, 2013; and D, 2014.

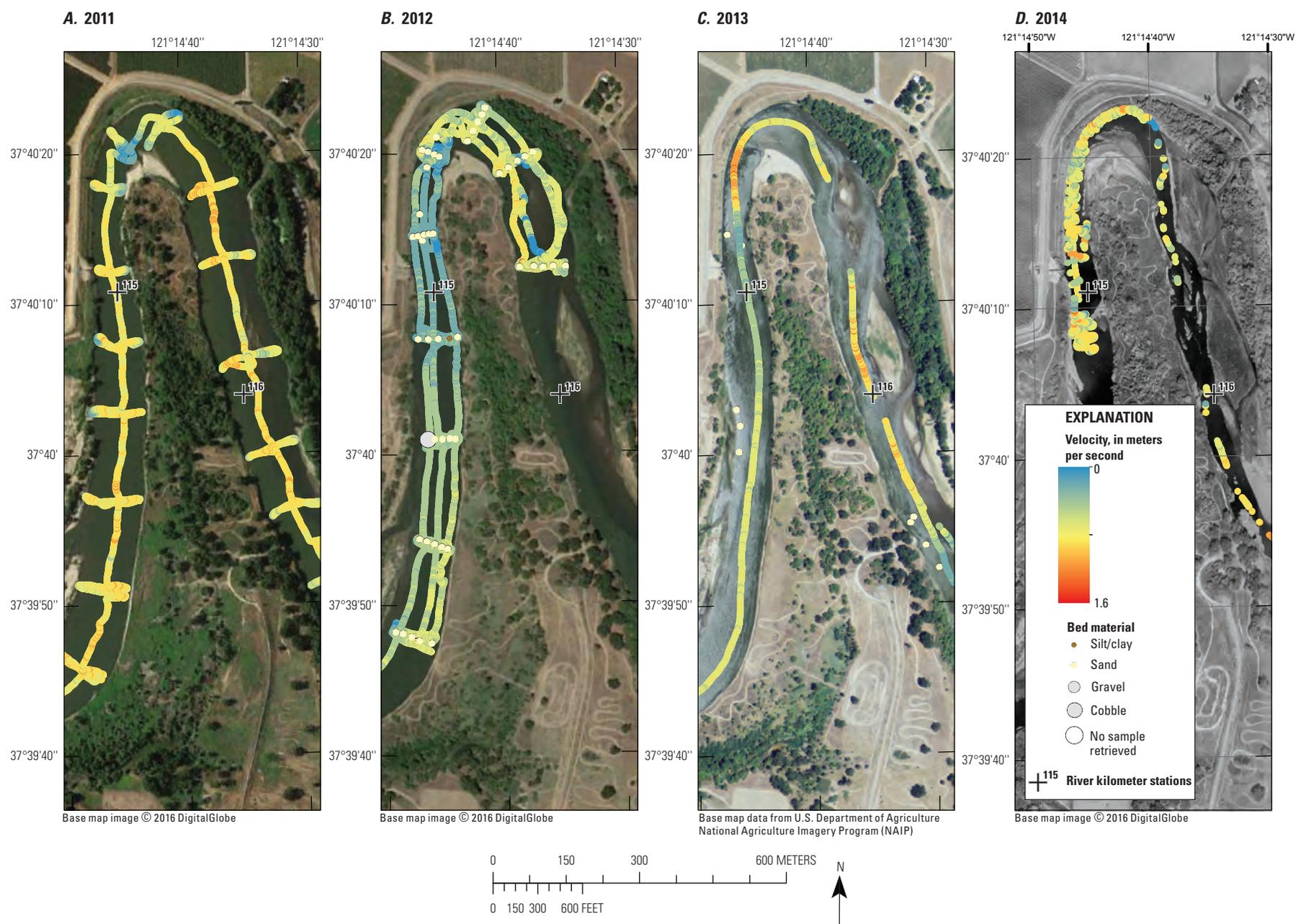


Figure 19. Streamflow depth-averaged velocity and bed-material size class in the Sturgeon Bend subreach of the lower San Joaquin River, California (river kilometers 111.9–116.3): *A*, 2011; *B*, 2012; *C*, 2013; and *D*, 2014.

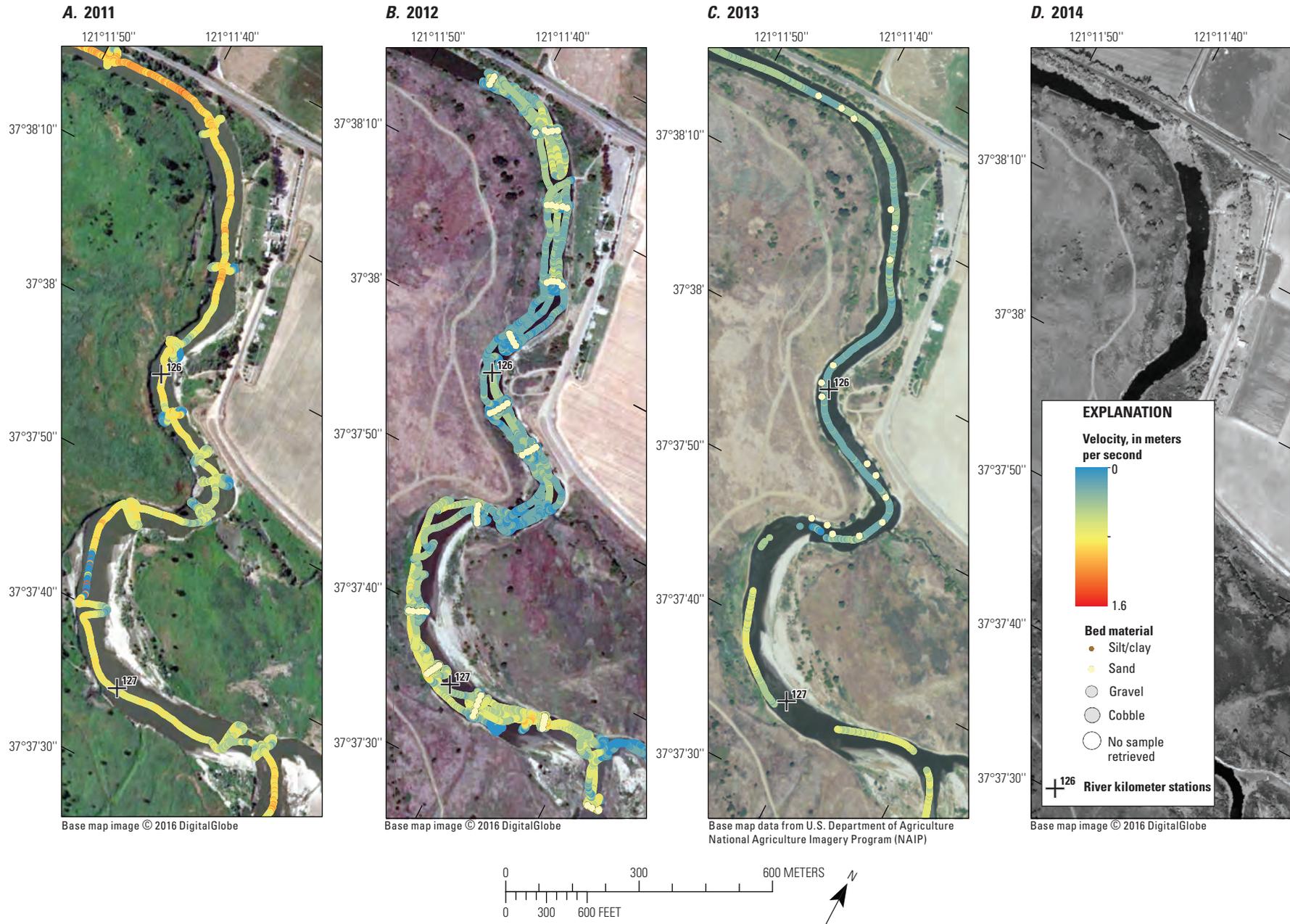


Figure 20. Streamflow depth-averaged velocity and bed-material size class in the Old Fisherman's Club subreach of the lower San Joaquin River, California (river kilometers 125.2–127.5): A, 2011; B, 2012; C, 2013; and D, 2014. No depth data or bed-material samples were collected in 2014 from this subreach.

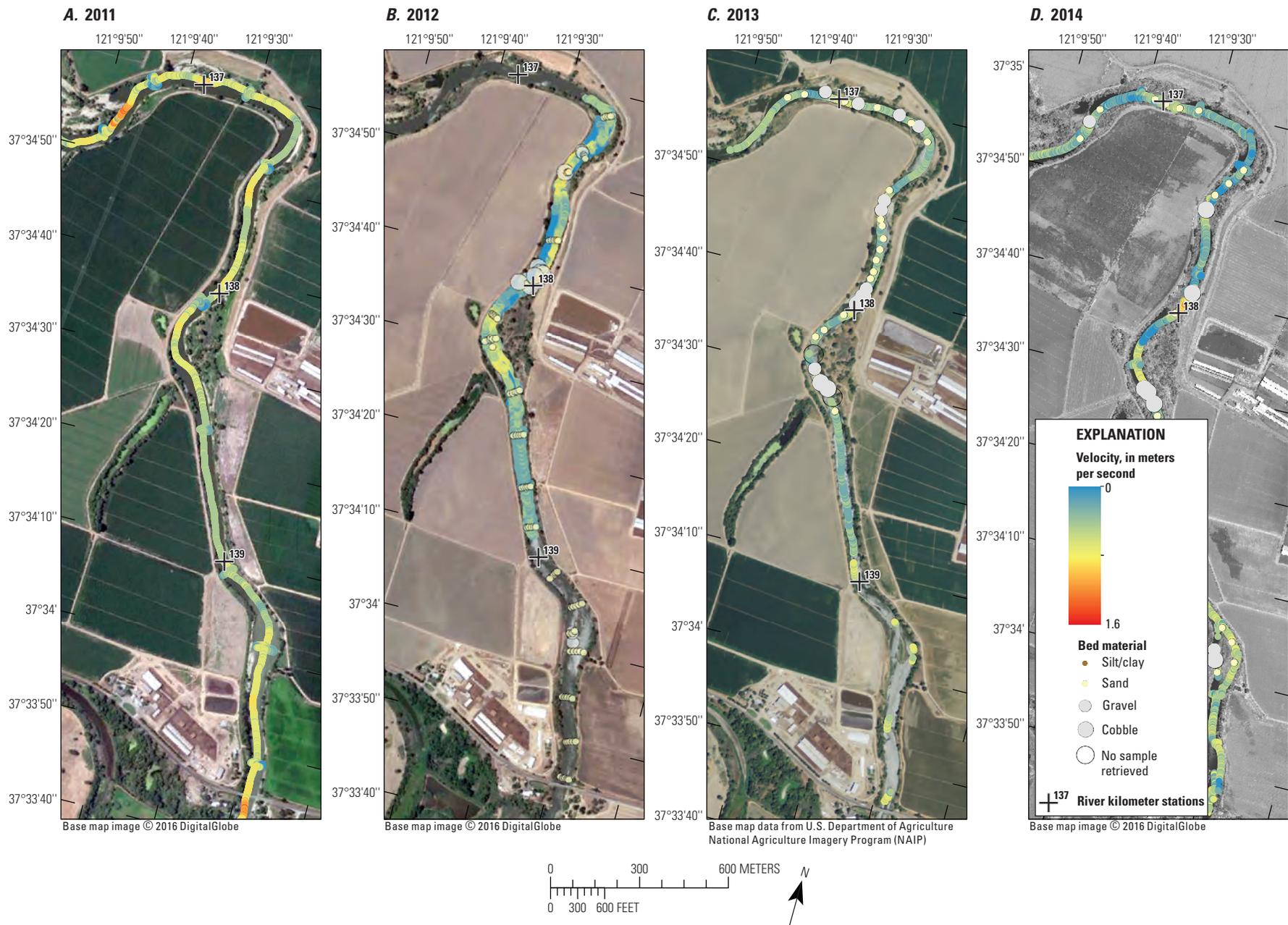


Figure 21. Streamflow depth-averaged velocity and bed-material size class in the Grayson Bridge subreach of the lower San Joaquin River, California (river kilometers 137.3–139.9): *A*, 2011; *B*, 2012; *C*, 2013; and *D*, 2014.



Figure 22. Cobble in the dry bed of Orestimba Creek at Eastin Road (river kilometer 12.8), approximately 5 kilometers northwest of Newman, California, April 24, 2014. The scale in the photograph is about 0.3 meter (1 foot) in length.



Figure 23. Edge of a large gravel bar in the lower San Joaquin River at river kilometer 146.3, the mouth of Del Puerto Creek, California, November 8, 2013. The scale in the photograph is about 1.2 meters (4 feet) in length.



Figure 24. Mouth of Del Puerto Creek, California, November 8, 2013, where the channel cut through a large gravel deposit.



Figure 25. Dry channel bed of Del Puerto Creek, 22 kilometers upstream from its confluence with the lower San Joaquin River, California, November 1, 2013. Gravel- and cobble-size sediment deposits were abundant in the creek bed.



Figure 26. Del Puerto Creek at Rogers Road (8.7 kilometers upstream from its confluence with the lower San Joaquin River, California), November 1, 2013. Deposits of gravel-size sediment were in the channel.



Figure 27. Exposed gravel bar in the lower San Joaquin River, California, near river kilometer 171.8, the mouth of Orestimba Creek, May 13, 2014. The exposed portion is about 4 meters across.

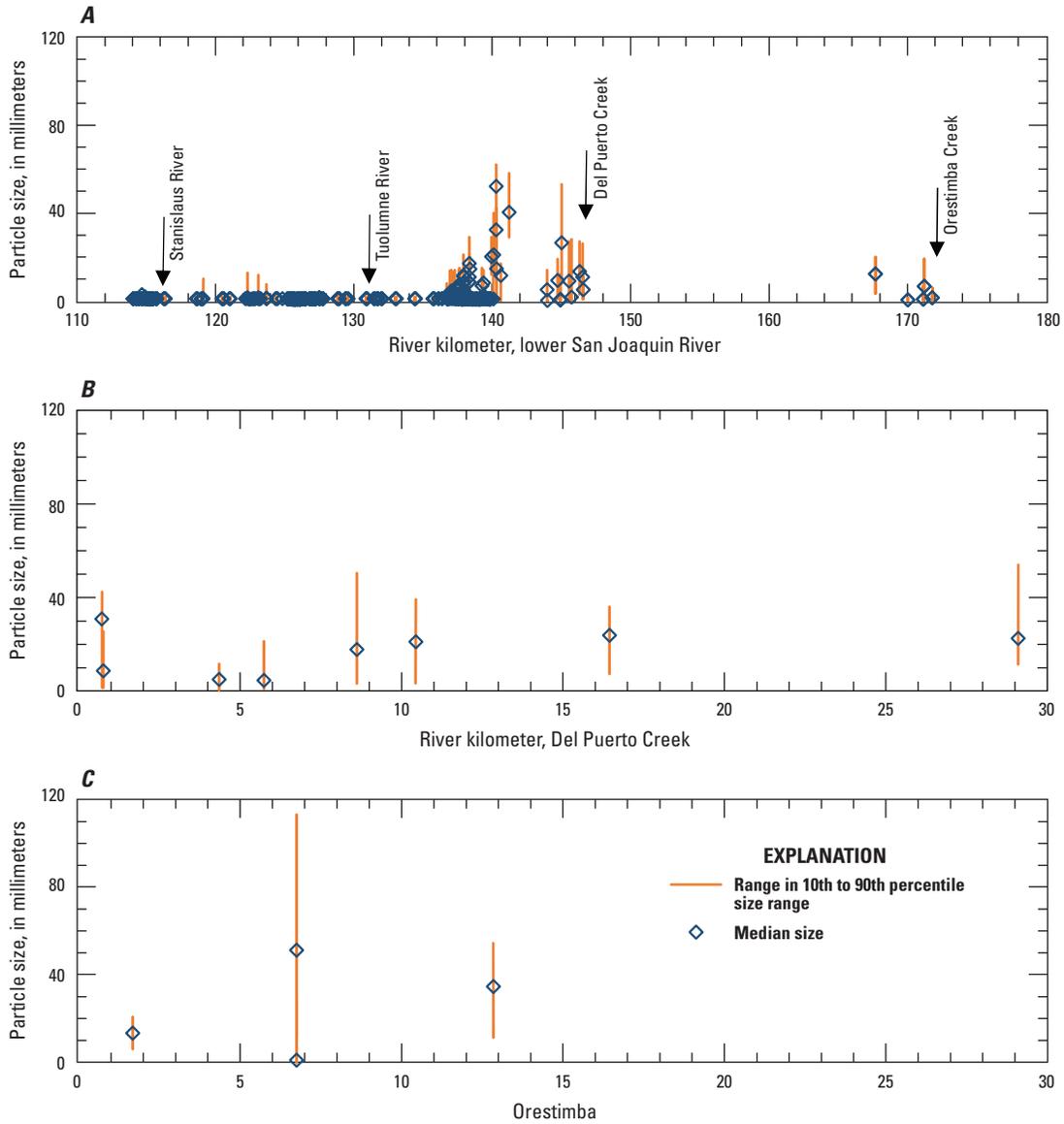


Figure 28. Particle-size distribution of bed-material samples collected from channels during 2011–13: *A*, lower San Joaquin River; *B*, Del Puerto Creek (enters the San Joaquin River at river kilometer 146.3); and *C*, Orestimba Creek (enters the San Joaquin River at river kilometer 171.8).

The dominant bed-material size class in the study areas of the lower San Joaquin River was sand. The Grayson Bridge subreach and an area immediately upstream from the Grayson Road Bridge were the only areas with large quantities of gravel in the streambed. Several exposed gravel bars were observed in this subreach, and gravel-sized bed material was typically found in the channel (as evidenced by bed-material samples) adjacent to those gravel bars. The areas of coarse (gravel-size) bed material in this subreach were separated by larger areas of sand.

Some small lenses of cobble-size sediment were in the banks next to Laird Park (Rkm 139.9–142.0); however, the bank material was not systematically sampled. The lenses of cobble-size sediment near Laird Park were generally less than 1 meter below the top of the bank and, in at least one location, appeared to come from an eroding bank and were mixed with man-made debris (bricks and pieces of metal), the likely remnants of an old levee berm. Because only small quantities of cobble (and no gravel) were observed in the banks, they did not appear to be the likely primary source of the gravel and cobble in the riverbed.

Gravel-size bed material in the lower San Joaquin River appeared to be primarily coming from Del Puerto Creek and, to a lesser extent, Orestimba Creek. Gravel bars were observed in several sites from the Coast Ranges to the lower San Joaquin River along the channel of Del Puerto Creek (figs. 24–26). In the most downstream reach of Del Puerto Creek (approximately 0.8–1.6 km upstream from the confluence), gravel-size sediment was found to depths of at least 10 cm in the streambed. Some gravel-size bed material was also present in the lower San Joaquin River at a few areas downstream from the confluence with Orestimba Creek (figs. 27, 28). Previous reports described gravel in the lower San Joaquin River only at the confluences (W.P. Smith, Assistant Engineer, to Colonel G.H. Mendell, June 28, 1883, as quoted in Reed and others, 1891, p. 89; Mussetter Engineering, Inc., and others, 2000); however, these reports were probably based only on exposed gravel bars. Samples collected in this study showed the patches of gravel bed material in the lower San Joaquin River extend farther downstream from Del Puerto Creek than from Orestimba Creek. The results plotted by river kilometer in figure 28 show the linear extent of gravel and cobble bed material in samples along the lower San Joaquin River as well as the gradual decrease in gravel and cobble downstream from Del Puerto Creek, where the bed-material composition transitions to nearly all sand and finer material.

Stream Slope

Because of the relatively flat valley floor, rivers in the Central Valley tend to have a very low gradient and, thus, generally do not have bed shear stresses high enough to transport coarse bed material. As a result, rivers on the Central Valley floor are primarily sand bedded (Mussetter

Engineering, Inc., and others, 2000; Singer and Dunne, 2004; Singer, 2008). The longitudinal elevation profiles of the San Joaquin River, three major east-side tributary rivers (Stanislaus, Tuolumne, and Merced), and two major west-side tributary creeks (Del Puerto and Orestimba) are shown in figure 29A, C). The profiles show a generally concave-upward, longitudinal elevation profile, which is typical of rivers (Mackin, 1948). The locations of some of the major dams, by river kilometer, also are shown in the profiles.

The channel gradient, which was calculated from the longitudinal profiles, is shown (as smoothed) in figures 29B, 29D. The channel slopes in the upper watershed of all four rivers (Stanislaus, Tuolumne, Merced, and San Joaquin) varied greatly as they transitioned from the Sierra Nevada to the foothills and then to the Central Valley. The downstream part of each river (0–50 km) had a gentle slope, generally around 0.0003–0.0005 meters per meter (m/m). The two creeks (Del Puerto and Orestimba) had slopes of about 0.005–0.010 m/m, an order of magnitude steeper than the rivers and similar to that of the river reaches upstream from the dams. The steeper gradient of the creeks is due to the shorter distance from their headwaters in the Coast Ranges to the lower San Joaquin River compared to the distance along rivers from their headwaters in the Sierra Nevada. Consequently, the peak flows in steep west-side creeks exert bed shear stresses capable of transporting coarse bed material to the predominantly sand-bedded reaches of the lower San Joaquin River.

Bed Aggradation, 2011–14

Aggradation of bed material in the lower San Joaquin River was observed in many areas during the study period, including Sturgeon Bend (fig. 30), where bathymetric data were collected in all 4 years of this study. Following high flows in 2011, Sturgeon Bend had a 14-m deep pool, which was filled with sediment—as much as about 7–8 m of deposition—by 2014. At the same time, the sand bar on the inner bank of the bend also grew, so that by 2014, the river was confined to a narrow, 1–3-m deep channel near the outer bank of Sturgeon Bend. The arbitrary datum shown in Figures 30 and 31 corresponds roughly to elevation in meters using North American Vertical Datum of 1988. This datum was determined by adjusting the difference in bed elevation surveyed in 2012 and adjusting it corresponding to differences in stage at the nearest streamgage during other longitudinal profiles.

Another area where water depth was measured every year from 2011 to 2014 was the Grayson Bridge subreach (fig. 31). At this site, a sediment wave approximately 2.5–4.0-m in height, composed of sand, advanced downstream at a rate of about 100 meters per year (m/yr). The channel in this longitudinal section (Rkm 138.8–139.2) is about 35 m wide. An estimated 9,500 plus or minus 1,000 cubic meters (m³) of sediment was deposited each year between 2011 and 2014 in this area.

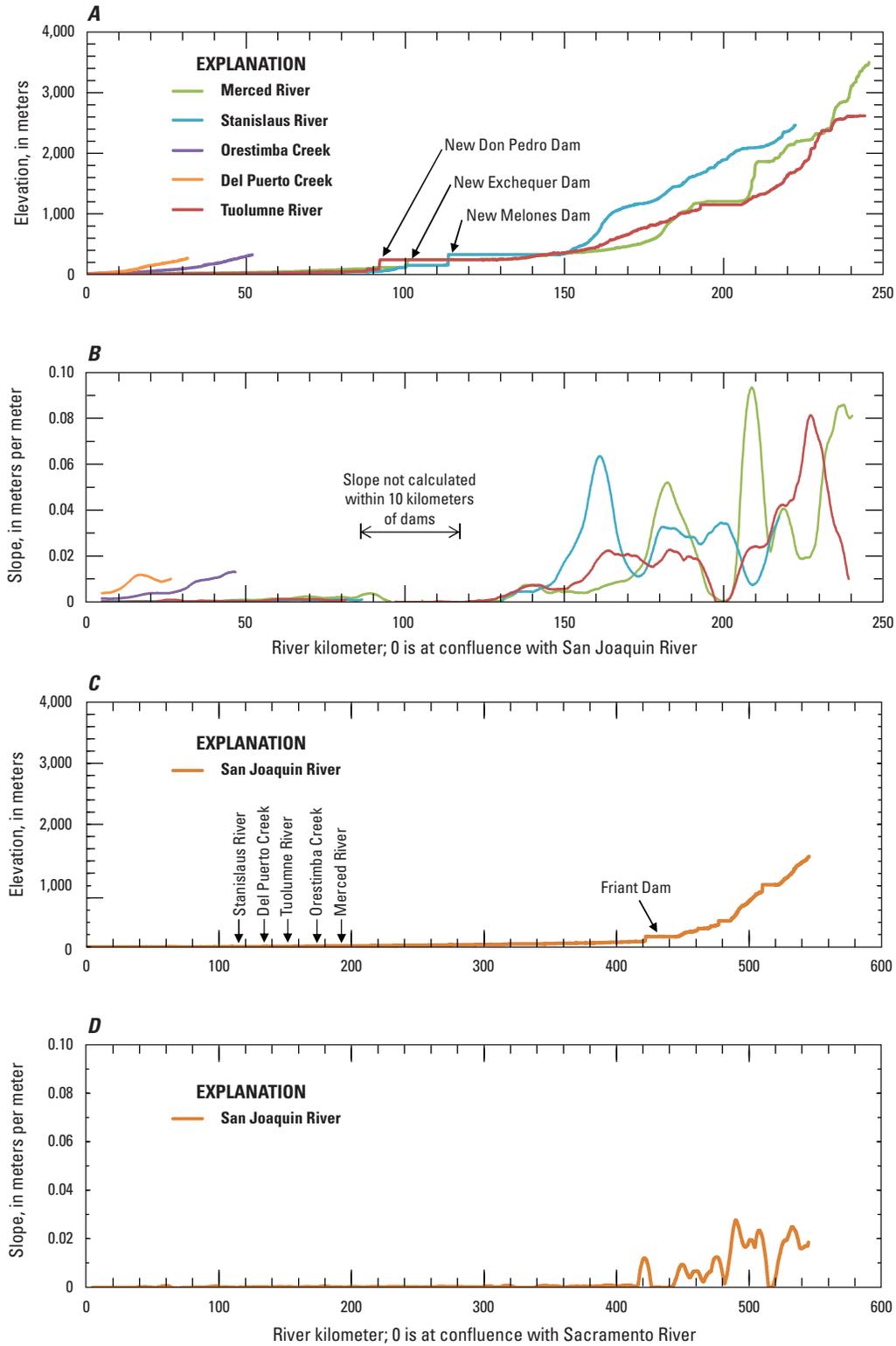


Figure 29. Longitudinal profiles of water-surface elevation, water-surface slope, and principal dams in rivers and creeks of the San Joaquin Basin, California: *A*, elevation of rivers and creeks tributary to the San Joaquin River; *B*, slope of rivers and creeks tributary to the San Joaquin River; *C*, elevation of the San Joaquin River; and *D*, slope of the San Joaquin River. The water-surface slope was smoothed using a 10-kilometer-averaging window.

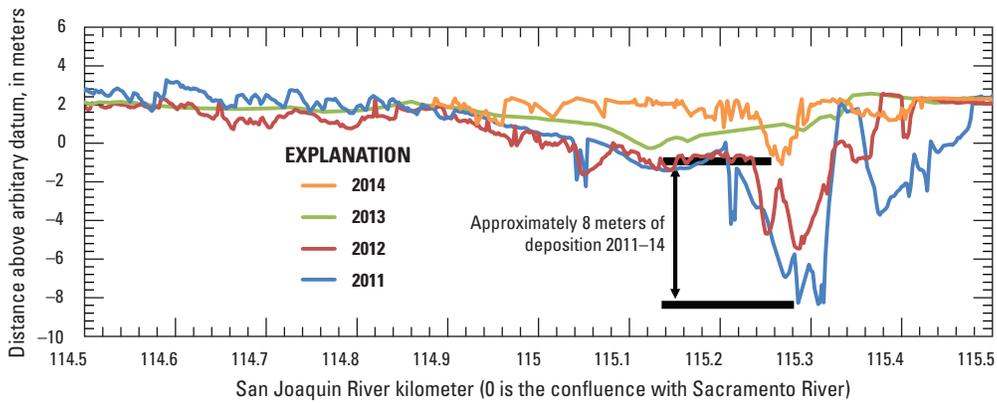


Figure 30. Bed-elevation changes along selected longitudinal section (river kilometers 114.5–115.5) in the Sturgeon Bend subreach, lower San Joaquin River, California, 2011–14. Longitudinal profile in 2013 followed the river centerline, which may not correspond to the deepest part (thalweg) of the river channel.

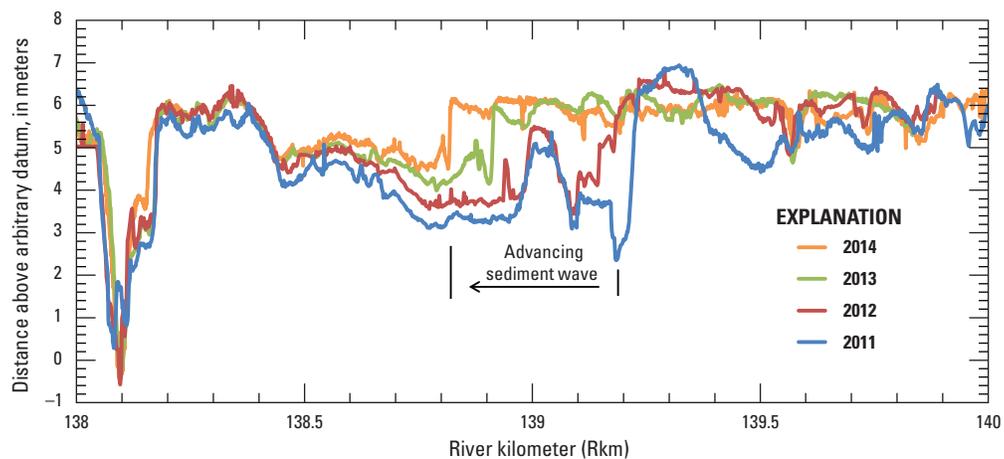


Figure 31. Bed-elevation changes along selected longitudinal section (river kilometers 138–140) in the Grayson Bridge subreach, San Joaquin River, California, 2011–14. Water depths measured each year; bed elevation was approximated from changes in stage at an upstream U.S. Geological Survey streamgage.

Discussion and Conclusions: Bed Material and Streamflow in the Lower San Joaquin River

The following section provides interpretation of the results related to (1) gravel in a sand-bedded river, (2) historical changes in hydrology, (3) aggradation of the river bed, and (4) sources of gravel and cobble.

Gravel in a Sand-Bedded River

The bed-material particle size in most rivers becomes finer downstream as a result of weathering and natural sorting concomitant with downstream decreases in slope; it often transitions between gravel and sand over a short distance (Smith and Ferguson, 1995; Knighton 1998; Ferguson, 2003). The presence of gravel in sand-bedded systems is often attributed to lateral inputs (such as the west-side tributary creeks in this study), but the drivers of coarse (gravel) bedload transport and the fate of such gravel in a sand-bedded river are not well understood. Most likely, the gravel moves slowly downstream, primarily during high flows, in the troughs of sand dunes until it reaches a low point in the river (Hassan and others, 1999; Venditti and others, 2010). Numerical sediment transport modeling can help estimate the sediment transport capacity of the of the lower San Joaquin River and address questions regarding the time required for gravel to be transported downstream from the Del Puerto reach.

Changes in Hydrology

We analyzed several years of peak-flow data and daily average stream discharge data (figs. 6, 7) and determined that the upstream dams reduced peak flows during most water-year types. Although runoff from the largest storms is not completely captured in reservoirs, the reservoir flood-control capacity seems to be sufficient to capture runoff from many of the small- to mid-size storms. The resulting effect is that average peak streamflow has been lower after dam construction for all water year types except for critical years, and the time between peak flows of similar magnitude has lengthened. The low peak flows during low-flow water years also could affect sediment transport and long-term sediment storage in the lower San Joaquin River.

Aggradation of the River Bed

Aggradation of sand in the lower San Joaquin River during the low-flow water years of this study could decrease

the availability of accessible gravel to sturgeon for spawning until subsequent high-flows scour the bed. Because sedimentation was recorded in the lower San Joaquin River even during dry years in some of areas in this study, and gravel migrates downstream through dune troughs until it reaches low points in the river, it is possible that gravel and cobble are in many sections of the river, but were buried and remained undetected during the 2012–14 bed-material sampling. If lenses of gravel and cobble are present beneath layers of sand, then high flows (such as those during 2011) could expose them, thereby increasing the sturgeon spawning habitat.

High stream discharge in July 2011 was associated with high water velocity (fig. 15); in water year 2011. This likely caused erosion which was followed by deposition during low flows. Deposition during the subsequent years indicates that sediment supply exceeds the transport capacity during those low-flow years. Hence, greater sediment-transport capacity during high-flows could evacuate sand from pools, access and potentially mobilize any underlying gravels, and cause deep scour holes, particularly at sharp meander bends (like Sturgeon Bend), as mobility thresholds for bed sediment are exceeded.

During low-flow years, the accumulation of sand-size sediment in the lower San Joaquin River could affect spawning habitat by filling in deep scour holes and by infiltrating the interstitial spaces or completely burying spawning gravels. Evaluation of sediment transport and bed evolution under a range of flow conditions, water-year types, and upstream reservoir flow-release scenarios could provide information about the conditions needed to establish and maintain sturgeon-breeding habitat in the lower San Joaquin River.

Sources of Gravel and Cobble

Bed material was sampled in many locations from 2012 to 2014. The bed of the lower San Joaquin River was predominately sand, except in the Del Puerto reach. In the Del Puerto reach, several exposed gravel bars and large patches of gravel were in the channel from the mouth of Del Puerto Creek and extending downstream approximately 10 km. Conversely, only a single exposed gravel bar was found at the mouth of Orestimba Creek, and only a few patches of gravel were found in the Orestimba Creek reach. It is possible that Orestimba Creek had delivered large quantities of gravel that were buried beneath sand at the time of this study. We did not evaluate the long-term stability of gravel bars, downstream-migration rates of gravel material, or the sediment-delivery rate of the tributary creeks. A sediment-transport study could evaluate these factors.

Several deep scour holes were measured in the 2011 survey, including Sturgeon Bend; however, no bed-material samples were collected that year. As a result, we could not determine whether coarse bed material in those areas was buried in later years. Given the distance from Del Puerto Creek, it seems likely that coarse bed material was buried in the channel before it could migrate downstream to some of the deep scour holes measured in 2011. In the major tributaries, the bed-material transitions to sand several kilometers before the confluence with the lower San Joaquin River. With dam-regulated streamflow regimes, it is unlikely that the larger tributaries are important sources of gravel or cobble sediment to the lower San Joaquin River.

Although gravel and cobble were found in the dry creek beds of Del Puerto and Orestimba Creeks, evidence indicated that Del Puerto Creek was the primary source of coarse bed material to the lower San Joaquin River and that Orestimba Creek did not appear to contribute much coarse bedload sediment.

Del Puerto Creek has a smaller watershed than Orestimba Creek, so it is not clear why Del Puerto Creek would deliver more coarse sediment to the lower San Joaquin River than Orestimba Creek. One possibility is that Orestimba Creek cannot transport as much coarse sediment as Del Puerto Creek because the channel slope is gentler in Orestimba Creek. The slope of the lower 10 km of Orestimba Creek was about 0.002 m/m, whereas the slope of Del Puerto Creek was about 0.004 m/m (fig. 29B). Another possibility is that the hydrologic and geomorphic history prior to the study period differed between the two watersheds. A third possibility is that peak flow near the mouth of the stream is much less than that near the mountains (fig. 6E). Orestimba Creek has been heavily modified by gravel mining and agricultural encroachment, which have changed the channel geometry and altered the natural sediment-transport regime (United States Army Corps of Engineers, 2013). During floods, streamflow can spill over the creek banks where the channel reaches the San Joaquin Valley floor and spread onto the floodplains (United States Army Corps of Engineers, 2013), which could reduce the transport capacity of the channel. As the primary, modern-day source of gravel and cobble bed material to the lower San Joaquin River, Del Puerto Creek appears to be important for the spawning success of sturgeon by supplying gravel for spawning habitat.

Summary

Decreases in the San Francisco Bay estuary white sturgeon population have led to an increased effort to understand their migration behavior and habitat preferences. Coarse bed material (gravel and cobbles), in particular, is important for the survival of white sturgeon eggs and larvae. Evidence of spawning by white sturgeon (*Acipenser transmontanus*) in the lower San Joaquin River has recently

been documented (2011–12) by the U.S. Fish and Wildlife Service. Knowledge of the physical characteristics of the lower San Joaquin River can be used to preserve sturgeon spawning habitat and lead to management decisions that could help increase the San Francisco Bay estuary white sturgeon population.

In order to characterize the hydraulic conditions and bed composition of the lower San Joaquin River reaches where sturgeon spawned, the U.S. Geological Survey (USGS) initiated scientific investigations, in cooperation with the U.S. Fish and Wildlife Service (USFWS), of selected reaches and tributaries of the lower river. This report documents the methods and findings of the 4-year study by (1) describing selected spawning reaches in terms of habitat-related physical characteristics (such as water depth and velocity, channel slope, and bed material) of the lower San Joaquin River between its confluences with the Stanislaus and Merced Rivers; (2) describing variations in these physical characteristics during wet and dry years; and (3) identifying potential reasons for these variations.

During high-flow periods, a substantial part of the San Joaquin River streamflow is diverted at the San Joaquin River Control Structure, but the diverted water eventually returns to the San Joaquin River through the Mariposa Bypass, where the lower San Joaquin River is considered to begin, 234 river kilometers (Rkm 234) upstream from its confluence with the Sacramento River. The lower river is a low-gradient, single-channel, meandering river. Streamflow in the lower San Joaquin River is primarily derived from the middle San Joaquin River (through the Mariposa Bypass) and from three tributary rivers flowing west from the Sierra Nevada, all regulated by dams. In contrast, tributary creeks flowing east from the Coast Ranges generally are not dammed and are ephemeral. Streamflow in the lower river is also affected by numerous agricultural diversions.

The lower San Joaquin River was divided into five study reaches, named for the tributary confluences at their upstream ends. In upstream order, the study reaches were the Stanislaus, Tuolumne, Del Puerto, Orestimba, and the Merced. Although data were collected from all study reaches, of special interest were three subreaches where the USFWS collected viable eggs at multiple sites they sampled in 2011–12. In upstream order, the subreaches were Sturgeon Bend (Rkm 114.0–116.3), Old Fisherman's Club (Rkm 125.2–127.5), and Grayson Bridge (Rkm 137.3–139.9). The USGS collected data for 3 to 4 years in those three subreaches, whereas in other reaches, data often were collected only for 1 or 2 years. The study data were collected during spring through summer of 4 water years during which runoff ranged from wet in 2011 and dry in 2012 to critical in 2013 and 2014. Data collection included approximately 105 channel cross sections and measured water depth and velocity in a total of 92 river kilometers by collecting nearly 150,000 depth-averaged water velocity measurements and 246,000 water-depth measurements.

Water depth was measured by soundings using a multibeam sonar and an ADCP. Depth and velocity data were collected using two different approaches: channel cross sections and longitudinal profiles. Bed forms and texture were imaged using side-scan sonar. Longitudinal elevation profiles of the San Joaquin River, three major east-side tributary rivers (Stanislaus, Tuolumne, and Merced), and two west-side tributary creeks (Del Puerto and Orestimba) were constructed using elevation data from the National Elevation Dataset and centerlines from the National Hydrography Dataset. Bed material generally was sampled from boats using a rotary-scoop bed-material sampler; thus, the surface layer of the bed was sampled, nominally to a depth of 5 cm. Particle-size distribution of the bed material was determined using a combination of sieving (five mesh sizes for the fraction less than 2 mm) and direct manual measurement. That is, particles larger than 2 mm were sorted manually and sized using a template (SAH-97 analyzer). Particles were classified using a generalized Wentworth scale as fine sediment (smaller than 0.063 mm, silt and clay), sand (0.063 to less than 2 mm), gravel (2 to less than 64 mm), and cobble (64 to less than 256 mm).

The bed of the lower San Joaquin River was predominately sand, but had areas of gravel. The hydrology of the lower San Joaquin River has been altered by the construction of several major dams on its tributary rivers. How these alterations in hydrology have affected long-term sediment-transport dynamics in the lower San Joaquin River was not determined, but this study documented aggradation in three subreaches during the 3 years following 2011 (a wet year). The aggradation of fine sediment during low-flow water years could reduce the availability of clean, exposed, coarse bed material for sturgeon to use during spawning. Coarse bed material was found in patches separated by large areas of sand. The gravel is likely to migrate slowly downstream, primarily during high flows, in the troughs of sand dunes until it reaches a low point in the river. Additional work would be required to quantify the supply of coarse sediment, estimate transport rates, and determine the fate of coarse sediment in the lower San Joaquin River.

The principal source of gravel and cobble was determined to be Del Puerto Creek and, to a lesser degree, Orestimba Creek (on the basis of locations of coarse material in the bed of the lower San Joaquin River, the proximity to the confluences with these two creeks, and presence of coarse bed material in the two creeks). Because Del Puerto Creek was the primary source of coarse sediment to the lower San Joaquin River in this study, it is important to the long-term spawning success of sturgeon in the lower San Joaquin River.

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Appendix 1: River Kilometer Station Locations

Table 1–1. River kilometer stations, and their associated geographic coordinates, based on the 2012 National Aerial Imagery Program (NAIP) aerial imagery of the lower San Joaquin River.

[River stationing based on channel configuration during 2012 NAIP aerial imagery. North American Datum of 1983. Easting and northing are map grid coordinates based on Universal Transverse Mercator, 10N, projection. **Abbreviations:** dd, decimal degrees; Rkm, river kilometer]

Rkm	Latitude (dd)	Longitude (dd)	Northing (meters)	Easting (meters)	Rkm	Latitude (dd)	Longitude (dd)	Northing (meters)	Easting (meters)
0	38.0646	-121.8565	4213596	600312	38	38.0664	-121.5571	4214166	626575
1	38.0624	-121.8459	4213367	601238	39	38.0589	-121.5546	4213336	626806
2	38.0543	-121.8417	4212470	601626	40	38.0571	-121.5434	4213149	627788
3	38.0454	-121.8398	4211490	601806	41	38.0557	-121.5322	4213011	628778
4	38.0370	-121.8356	4210564	602177	42	38.0543	-121.5209	4212872	629768
5	38.0296	-121.8292	4209746	602751	43	38.0527	-121.5099	4212709	630739
6	38.0237	-121.8206	4209106	603514	44	38.0460	-121.5027	4211979	631384
7	38.0199	-121.8106	4208690	604402	45	38.0404	-121.4939	4211365	632164
8	38.0207	-121.7994	4208794	605378	46	38.0348	-121.4849	4210765	632960
9	38.0187	-121.7885	4208586	606346	47	38.0295	-121.4758	4210183	633773
10	38.0194	-121.7771	4208670	607338	48	38.0245	-121.4663	4209642	634612
11	38.0229	-121.7667	4209077	608247	49	38.0169	-121.4610	4208804	635095
12	38.0252	-121.7560	4209346	609188	50	38.0091	-121.4553	4207949	635612
13	38.0256	-121.7446	4209395	610186	51	38.0013	-121.4496	4207088	636120
14	38.0283	-121.7338	4209706	611128	52	37.9958	-121.4408	4206496	636901
15	38.0345	-121.7256	4210410	611835	53	37.9943	-121.4305	4206342	637817
16	38.0387	-121.7158	4210892	612695	54	37.9955	-121.4193	4206498	638797
17	38.0419	-121.7051	4211256	613626	55	37.9926	-121.4087	4206186	639734
18	38.0469	-121.6958	4211826	614438	56	37.9889	-121.3983	4205788	640648
19	38.0546	-121.6899	4212688	614939	57	37.9841	-121.3887	4205274	641499
20	38.0607	-121.6816	4213371	615661	58	37.9773	-121.3812	4204537	642171
21	38.0693	-121.6788	4214329	615890	59	37.9707	-121.3736	4203811	642858
22	38.0782	-121.6804	4215316	615735	60	37.9640	-121.3659	4203085	643544
23	38.0871	-121.6799	4216296	615763	61	37.9582	-121.3575	4202446	644291
24	38.0939	-121.6730	4217068	616364	62	37.9546	-121.3471	4202071	645215
25	38.0962	-121.6621	4217333	617315	63	37.9518	-121.3363	4201769	646165
26	38.0928	-121.6517	4216974	618233	64	37.9451	-121.3415	4201022	645725
27	38.0882	-121.6422	4216472	619069	65	37.9390	-121.3426	4200338	645643
28	38.0898	-121.6315	4216662	620010	66	37.9358	-121.3332	4200003	646476
29	38.0978	-121.6283	4217550	620277	67	37.9293	-121.3278	4199290	646960
30	38.1024	-121.6194	4218072	621046	68	37.9219	-121.3220	4198475	647488
31	38.1039	-121.6082	4218260	622028	69	37.9150	-121.3224	4197711	647461
32	38.1022	-121.5971	4218088	622999	70	37.9068	-121.3246	4196794	647290
33	38.0993	-121.5864	4217781	623949	71	37.8990	-121.3265	4195926	647138
34	38.0961	-121.5761	4217434	624856	72	37.8908	-121.3286	4195011	646968
35	38.0873	-121.5737	4216466	625081	73	37.8825	-121.3320	4194091	646685
36	38.0786	-121.5722	4215499	625223	74	37.8739	-121.3323	4193140	646679
37	38.0733	-121.5631	4214921	626038	75	37.8663	-121.3274	4192298	647120

Table 1–1. River kilometer stations, and their associated geographic coordinates, based on the 2012 National Aerial Imagery Program (NAIP) aerial imagery of the lower San Joaquin River.—Continued

[River stationing based on channel configuration during 2012 NAIP aerial imagery. North American Datum of 1983. Easting and northing are map grid coordinates based on Universal Transverse Mercator, 10N, projection. **Abbreviations:** dd, decimal degrees; Rkm, river kilometer]

Rkm	Latitude (dd)	Longitude (dd)	Northing (meters)	Easting (meters)	Rkm	Latitude (dd)	Longitude (dd)	Northing (meters)	Easting (meters)
76	37.8610	-121.3204	4191717	647751	117	37.6593	-121.2406	4169471	655186
77	37.8540	-121.3202	4190947	647783	118	37.6526	-121.2349	4168736	655706
78	37.8496	-121.3236	4190448	647489	119	37.6499	-121.2243	4168456	656645
79	37.8420	-121.3202	4189616	647805	120	37.6429	-121.2283	4167671	656312
80	37.8346	-121.3153	4188804	648245	121	37.6381	-121.2227	4167150	656810
81	37.8269	-121.3100	4187959	648727	122	37.6341	-121.2147	4166717	657522
82	37.8220	-121.3181	4187393	648027	123	37.6273	-121.2111	4165970	657858
83	37.8175	-121.3149	4186905	648319	124	37.6347	-121.2080	4166794	658121
84	37.8127	-121.3175	4186362	648099	125	37.6378	-121.1982	4167157	658976
85	37.8089	-121.3273	4185929	647246	126	37.6326	-121.1923	4166591	659502
86	37.8060	-121.3185	4185627	648024	127	37.6266	-121.1897	4165929	659749
87	37.8001	-121.3128	4184975	648535	128	37.6214	-121.1837	4165359	660288
88	37.7927	-121.3122	4184156	648606	129	37.6132	-121.1816	4164457	660494
89	37.7913	-121.3073	4184006	649042	130	37.6109	-121.1719	4164219	661356
90	37.7837	-121.3038	4183170	649366	131	37.6060	-121.1782	4163663	660807
91	37.7760	-121.3022	4182316	649520	132	37.5988	-121.1792	4162862	660732
92	37.7690	-121.3035	4181536	649419	133	37.5991	-121.1703	4162905	661523
93	37.7622	-121.3090	4180774	648950	134	37.5951	-121.1772	4162458	660918
94	37.7552	-121.3047	4180003	649342	135	37.5894	-121.1719	4161831	661403
95	37.7512	-121.2977	4179573	649968	136	37.5835	-121.1676	4161182	661792
96	37.7461	-121.2911	4179022	650557	137	37.5832	-121.1605	4161164	662423
97	37.7388	-121.2973	4178196	650026	138	37.5771	-121.1580	4160494	662652
98	37.7301	-121.2975	4177236	650024	139	37.5692	-121.1554	4159623	662902
99	37.7231	-121.2931	4176461	650426	140	37.5614	-121.1524	4158756	663180
100	37.7150	-121.2890	4175569	650807	141	37.5548	-121.1487	4158029	663522
101	37.7107	-121.2798	4175111	651624	142	37.5598	-121.1460	4158595	663751
102	37.7148	-121.2770	4175574	651865	143	37.5595	-121.1361	4158576	664627
103	37.7204	-121.2720	4176195	652292	144	37.5563	-121.1279	4158234	665358
104	37.7206	-121.2611	4176235	653254	145	37.5513	-121.1226	4157688	665842
105	37.7139	-121.2673	4175487	652722	146	37.5441	-121.1183	4156896	666234
106	37.7058	-121.2720	4174582	652326	147	37.5377	-121.1141	4156193	666619
107	37.6982	-121.2730	4173733	652253	148	37.5329	-121.1122	4155661	666795
108	37.6907	-121.2732	4172905	652249	149	37.5279	-121.1106	4155118	666953
109	37.6870	-121.2698	4172502	652558	150	37.5257	-121.1023	4154880	667687
110	37.6876	-121.2592	4172580	653488	151	37.5205	-121.1001	4154307	667896
111	37.6817	-121.2610	4171924	653345	152	37.5136	-121.0956	4153549	668308
112	37.6745	-121.2646	4171118	653039	153	37.5082	-121.0883	4152973	668963
113	37.6691	-121.2557	4170532	653833	154	37.5010	-121.0842	4152180	669340
114	37.6622	-121.2492	4169776	654424	155	37.4929	-121.0799	4151282	669745
115	37.6697	-121.2460	4170614	654689	156	37.4880	-121.0706	4150761	670577
116	37.6678	-121.2431	4170407	654947	157	37.4811	-121.0702	4149997	670628

Table 1–1. River kilometer stations, and their associated geographic coordinates, based on the 2012 National Aerial Imagery Program (NAIP) aerial imagery of the lower San Joaquin River.—Continued

[River stationing based on channel configuration during 2012 NAIP aerial imagery. North American Datum of 1983. Easting and northing are map grid coordinates based on Universal Transverse Mercator, 10N, projection. **Abbreviations:** dd, decimal degrees; Rkm, river kilometer]

Rkm	Latitude (dd)	Longitude (dd)	Northing (meters)	Easting (meters)
158	37.4830	-121.0639	4150214	671183
159	37.4784	-121.0559	4149714	671900
160	37.4736	-121.0483	4149197	672578
161	37.4691	-121.0434	4148710	673028
162	37.4645	-121.0350	4148221	673780
163	37.4604	-121.0395	4147753	673391
164	37.4549	-121.0402	4147147	673344
165	37.4496	-121.0345	4146560	673856
166	37.4464	-121.0262	4146223	674595
167	37.4418	-121.0232	4145715	674870
168	37.4340	-121.0228	4144858	674924
169	37.4311	-121.0133	4144554	675775
170	37.4259	-121.0095	4143982	676127
171	37.4285	-121.0004	4144281	676921
172	37.4209	-121.0022	4143435	676778
173	37.4150	-121.0023	4142788	676783
174	37.4113	-120.9963	4142389	677323
175	37.4057	-120.9967	4141759	677302
176	37.4032	-120.9903	4141493	677874
177	37.3984	-120.9841	4140980	678437
178	37.3924	-120.9913	4140299	677813
179	37.3865	-120.9944	4139639	677556
180	37.3868	-120.9864	4139684	678262
181	37.3810	-120.9843	4139044	678457
182	37.3760	-120.9816	4138495	678711
183	37.3757	-120.9737	4138479	679410
184	37.3701	-120.9779	4137844	679051
185	37.3674	-120.9767	4137553	679169
186	37.3628	-120.9788	4137035	678990
187	37.3542	-120.9770	4136087	679173

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