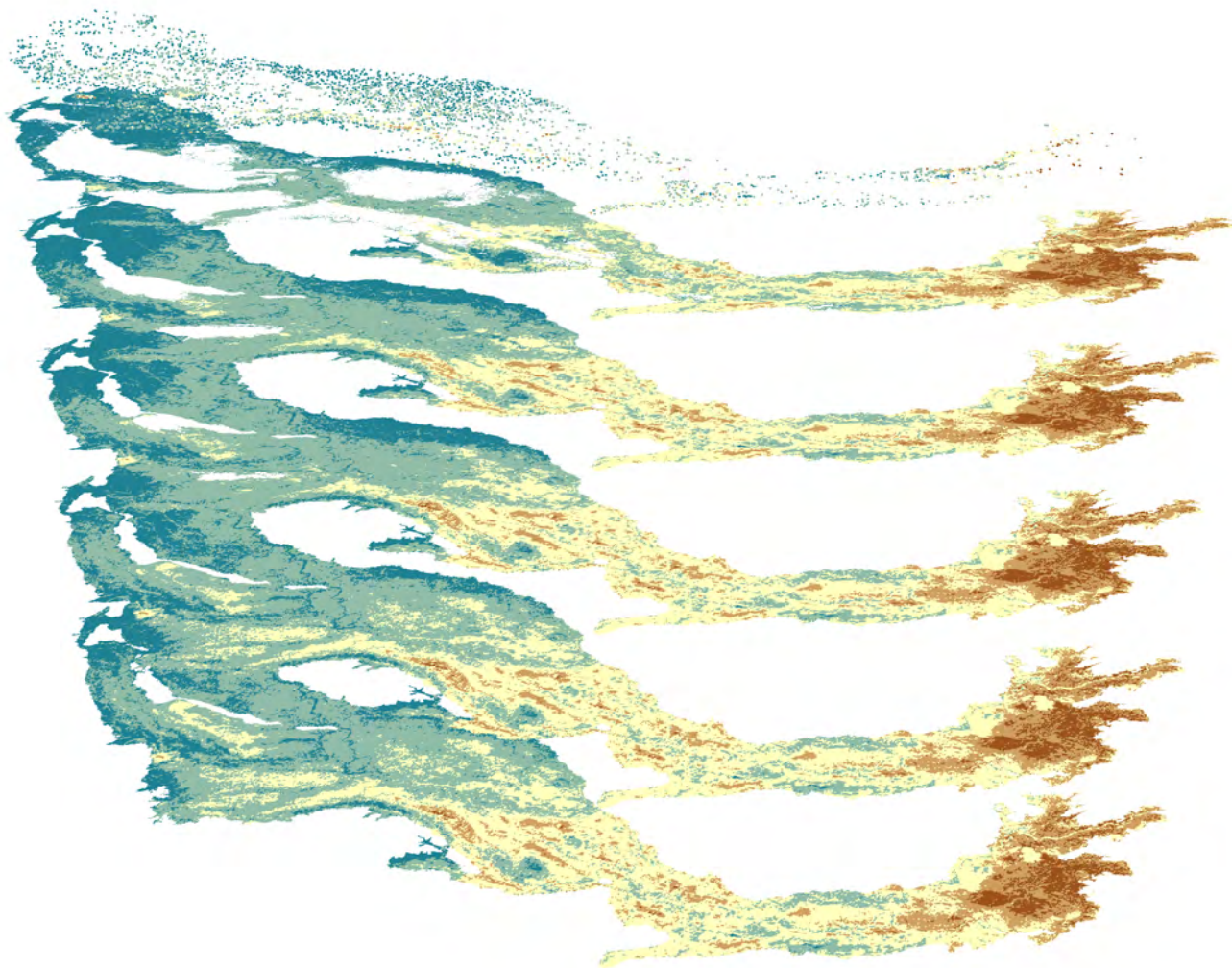


Water Availability and Use Science Program

Machine-Learning Predictions of Groundwater Specific Conductance in the Mississippi Alluvial Plain, South-Central United States, With Evaluation of Regional Geophysical Aerial Electromagnetic Data as Explanatory Variables



Scientific Investigations Report 2023–5099

Machine-Learning Predictions of Groundwater Specific Conductance in the Mississippi Alluvial Plain, South-Central United States, With Evaluation of Regional Geophysical Aerial Electromagnetic Data as Explanatory Variables

By Courtney D. Killian and Katherine J. Knierim

Water Availability and Use Science Program

Scientific Investigations Report 2023–5099

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[<https://doi.org/10.3133/sir20235099>]

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Conversion Factors

U.S. customary units to International System of Units

Multiply	By	To obtain
Length		
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
Area		
square mile (mi ²)	2.590	square kilometer (km ²)
Flow rate		
million gallons per day (Mgal/d)	0.04381	cubic meter per second (m ³ /s)

International System of Units to U.S. customary units

Multiply	By	To obtain
Length		
centimeter (cm)	0.3937	inch (in.)
millimeter (mm)	0.03937	inch (in.)
meter (m)	3.281	foot (ft)
kilometer (km)	0.6214	mile (mi)
Volume		
liter (L)	0.2642	gallon (gal)

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as follows:

$$^{\circ}\text{F} = (1.8 \times ^{\circ}\text{C}) + 32.$$

Datum

Vertical coordinate information is referenced to the North American Vertical Datum of 1988 (NAVD 88).

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Altitude, as used in this report, refers to distance above the vertical datum.

Supplemental Information

Specific conductance is given in microsiemens per centimeter at 25 degrees Celsius ($\mu\text{S}/\text{cm}$ at 25 °C).

Concentrations of chemical constituents in water are given in either milligrams per liter (mg/L) or parts per million (ppm).

Abbreviations

>	greater than
<	less than
$\mu\text{S}/\text{cm}$	microsiemens per centimeter at 25 degrees Celsius
1SE	one standard error
AEM	aerial electromagnetic
BRT	boosted regression tree
CV	cross validation
EDM	empirical distribution matching
id	interaction depth
ML	machine learning
mres	average inverted resistivity
n	number of samples
nt	number of trees
PDP	partial dependence plot
ppm	part per million
R^2	coefficient of determination
RMSE	root-mean-square error
SC	specific conductance
sh	shrinkage
TDS	total dissolved solids

Machine-Learning Predictions of Groundwater Specific Conductance in the Mississippi Alluvial Plain, South-Central United States, With Evaluation of Regional Geophysical Aerial Electromagnetic Data as Explanatory Variables

By Courtney D. Killian and Katherine J. Knierim

Abstract

The Mississippi Alluvial Plain, located in the south-central United States, is undergoing long-term groundwater-level declines within the surficial Mississippi River Valley alluvial aquifer (hereinafter referred to as “alluvial aquifer”), which has raised concerns about future groundwater availability. In some parts of the alluvial aquifer, groundwater availability for common uses such as irrigation, public supply, and domestic use is limited by quality (for example, high salinity) rather than quantity of water stored in the aquifer. The Mississippi Alluvial Plain region has an abundance of water-quality measurements in the alluvial aquifer and deeper aquifers; however, large areas lack direct measurements of salinity to evaluate regional groundwater availability. Statistical models can interpolate between wells to fill in spatial data gaps. In 2021, the U.S. Geological Survey trained two boosted regression tree (BRT) machine-learning models on specific conductance data available between 1942 and 2020 to predict spatially continuous surfaces of groundwater salinity at multiple depths for the alluvial aquifer and deeper aquifers. Well construction information, water levels, and surficial variables such as geomorphology and soils were included as explanatory variables in this baseline model. Additionally, subsurface electrical resistivity data from the first aquifer-wide aerial electromagnetic (AEM) survey for the region were incorporated to create a geophysical model. This work expands on prior BRT salinity predictions of the alluvial aquifer and extends predictions south to the Gulf of Mexico, where groundwater salinity is high. AEM survey data were not available for the southern extent of the alluvial aquifer at the time of modeling. A BRT model was trained without (baseline) and with (geophysical) AEM variables to test the ability of the models to predict salinity where explanatory data are missing and response data are sparse. Additionally, model sensitivity to AEM survey data was evaluated to better understand how AEM variables influence specific conductance predictions. Model performance was improved with the addition of geophysical data, which added three-dimensional

information, thereby improving salinity predictions at depth. Groundwater specific conductance predictions can help inform other geophysical investigations in the southern extent of the study area, where high groundwater specific conductance can obfuscate changes in aquifer sediment resistivity and could limit groundwater resources for agricultural, public supply, and domestic uses.

Introduction

High groundwater salinity can have major environmental and economic impacts by limiting available groundwater resources. High salinity can have negative effects on agriculture by inhibiting crop growth and reducing crop yield, depending on the salinity threshold of the species (Maas, 1993; Hardke, 2021). Groundwater with high salinity that is pumped for irrigation and then flows into nearby streams may interfere with stream ecology (Hart and others, 1990). Additionally, high salinity impacts the amount of freshwater available for human consumption; for example, the U.S. Environmental Protection Agency has a secondary maximum contaminant level of 500 milligrams per liter (mg/L) total dissolved solids (TDS; a measure of salinity) in drinking water (U.S. Environmental Protection Agency, 2022).

Effects of high groundwater salinity are observable worldwide, and impacts are more perceptible in areas where groundwater is the main source of freshwater (Welch and Hanor, 2011; Krishan, 2019). The Mississippi Alluvial Plain ([fig. 1](#)), located in the south-central United States, is no exception. Much of the groundwater in the Mississippi Alluvial Plain region is a freshwater resource with TDS concentrations less than 1,000 parts per million (ppm) (Kingsbury and others, 2014; Stanton and others, 2017), but concentrations increase—to as great as approximately 35,000 ppm—toward the Gulf of Mexico (Grubb, 1998). Additionally, there are several major areas of known high salinity in the Mississippi Alluvial Plain region: around Iberville Parish, Louisiana (Welch and Hanor, 2011);

in southeastern Arkansas (Larsen and others, 2021); and around White County, Ark. (Bryant and others, 1985), as well as in smaller, localized areas that may limit the use of groundwater for agricultural, public drinking water, and domestic uses.

The Mississippi Alluvial Plain region contains the highly productive surficial Mississippi River Valley alluvial aquifer (hereinafter referred to as “alluvial aquifer”) (Ackerman, 1996; Barlow and Clark, 2011; Clark and others, 2011; Lovelace and others, 2020). The alluvial aquifer was ranked first among principal aquifers in the United States in 2015 in estimated groundwater withdrawals for irrigation, supplying about 11,745 million gallons per day (Lovelace and others, 2020). The alluvial aquifer supports an \$11.9 billion industrial-scale agricultural economy (Alhassan and others, 2019). Tertiary aquifers that underlie the alluvial aquifer supply freshwater for agricultural, public drinking water, and domestic uses (Lovelace and others, 2020). Groundwater withdrawals in the alluvial aquifer and deeper aquifers have altered flow paths and created large cones of depression and have been linked to a migrating freshwater-saltwater interface and subsidence in some areas (Grubb, 1998; Clark and Hart, 2009; Jones and others, 2016). Groundwater resources in the Mississippi Alluvial Plain region are at risk for declining availability (Stanton and others, 2017). Observed declines in groundwater levels and decreased discharge to streamflow over time have prompted investigations to quantify the available groundwater resources in the Mississippi Alluvial Plain region (Ackerman, 1996; Clark and others, 2011; Killian and others, 2019; Yasarer and others, 2020).

Sources of salinity in the Mississippi Alluvial Plain region include but are not limited to upwelling from deeper saline aquifers, especially along regional fault lines (Kresse and Clark, 2008; Schrader, 2010; Larsen and others, 2021); subsurface dissolution of salt diapirs (Welch and Hanor, 2011); and saltwater intrusion from the Gulf of Mexico (Williamson and others, 1990; Grubb, 1998). Additionally, salinity may be concentrated at the surface as a result of evaporation where infiltration rates are low (Kresse and Clark, 2008).

Several techniques and methods have been developed to map spatial changes in groundwater quality, including geostatistical interpolation (Arslan, 2012), aerial electromagnetic (AEM) surveys of subsurface electrical conductivity (Delsman and others, 2018), and machine learning (ML) (Knierim and others, 2020). Multivariate regression analyses of hydrogeologic factors that may drive TDS concentrations were used to predict the occurrence and probability of exceedance for brackish groundwater at the national scale but were not detailed enough to delineate areas of high salinity in the Mississippi Alluvial Plain region (Stanton and others, 2017). ML predictions of groundwater salinity exist for the Mississippi embayment; however, these predictions were shown at only a single depth within the alluvial aquifer (Knierim and others, 2020). Groundwater salinity predictions at multiple depths are needed to understand how salinity changes vertically within an aquifer.

Salinity is typically quantified with TDS concentrations, but specific conductance (SC) is an easily collected measurement that describes dissolved ions. SC measures the electrical conductivity of water (Hem, 1985; Miller and others, 1988) and can correlate with concentrations of chloride and TDS, both of which have drinking water standards (U.S. Environmental Protection Agency, 2022). The Mississippi Alluvial Plain region has extensive SC data that have been collected over many decades (fig. 2) by the U.S. Geological Survey in cooperation with State agencies, including the Arkansas Natural Resources Division, the Louisiana Department of Transportation and Development, and the Mississippi Department of Environmental Quality. Although the SC dataset across the Mississippi Alluvial Plain region is large, there are still areas that do not have SC measurements, and the variation of SC with depth in the alluvial aquifer has not been characterized across the whole system.

In 2021, the U.S. Geological Survey trained two boosted regression tree (BRT) ML models on SC data available between 1942 and 2020 to predict spatially continuous surfaces of groundwater salinity at multiple depths for the alluvial aquifer and deeper aquifers. This study uses ML to predict SC across the Mississippi Alluvial Plain region for the alluvial aquifer and deeper aquifers. ML can be used to predict groundwater quality across aquifers where data are unavailable by identifying patterns of explanatory variables in complex datasets (Shangguan and others, 2017; Ransom and others, 2022). A BRT model is a type of ML model that provides model stability, resistance to collinearity among predictors, and the ability to handle missing input data (Elith and others, 2008; Kuhn and Johnson, 2013). Previous studies have demonstrated that BRT models can produce accurate predictions of groundwater quality across aquifer systems (Ransom and others, 2017; Rosecrans and others, 2017; Knierim and others, 2020). This study uses BRT models to expand on efforts to model groundwater SC in the Mississippi embayment (Knierim and others, 2020), extending the predictions to the Gulf of Mexico and adding high-resolution predictions for multiple depths. Response and explanatory variables and model predictions are available in the companion U.S. Geological Survey data release (Killian and Knierim, 2023).

The objectives of this study were to (1) develop BRT models to predict SC across the Mississippi Alluvial Plain region at multiple depths for the alluvial aquifer and deeper aquifers, (2) test the ability of the models to predict where there is a paucity of groundwater SC data and explanatory variables, and (3) evaluate model performance and sensitivity to the incorporation of explanatory variables from a regional geophysical survey. Model performance was assessed by generating two BRT models, a “baseline” model that did not incorporate geophysical data and a “geophysical” model that did. The predictions will help identify where high SC could limit groundwater resources for agricultural, public drinking water, and domestic uses. The predictions will also aid in the interpretation of geophysical data by identifying where high SC may obfuscate variations in aquifer sediment resistivity.

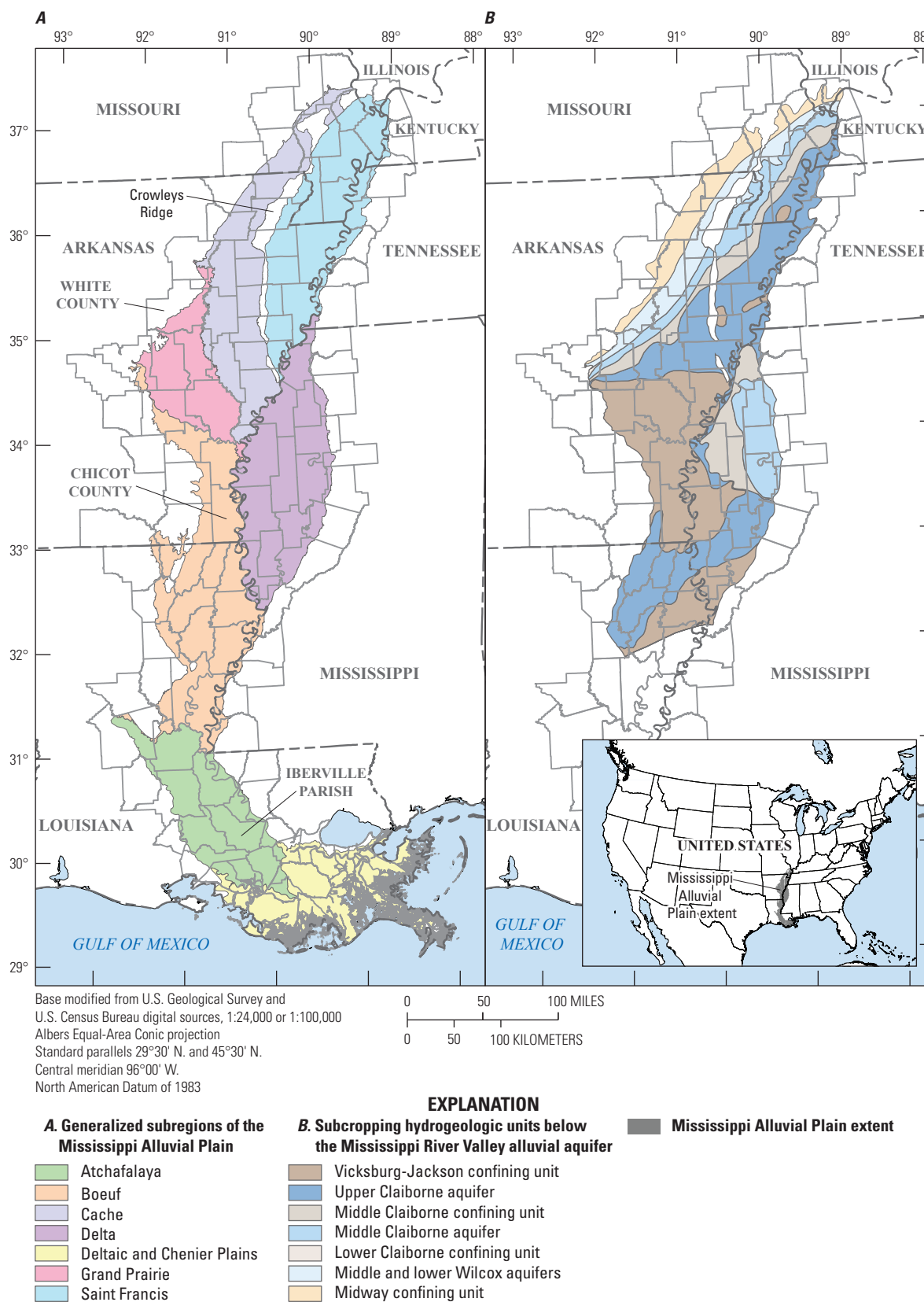


Figure 1. Overview of the Mississippi Alluvial Plain, south-central United States, by **A**, generalized subregions (Ladd and Travers, 2019) and **B**, subcropping hydrogeologic units (Hart and others, 2008).

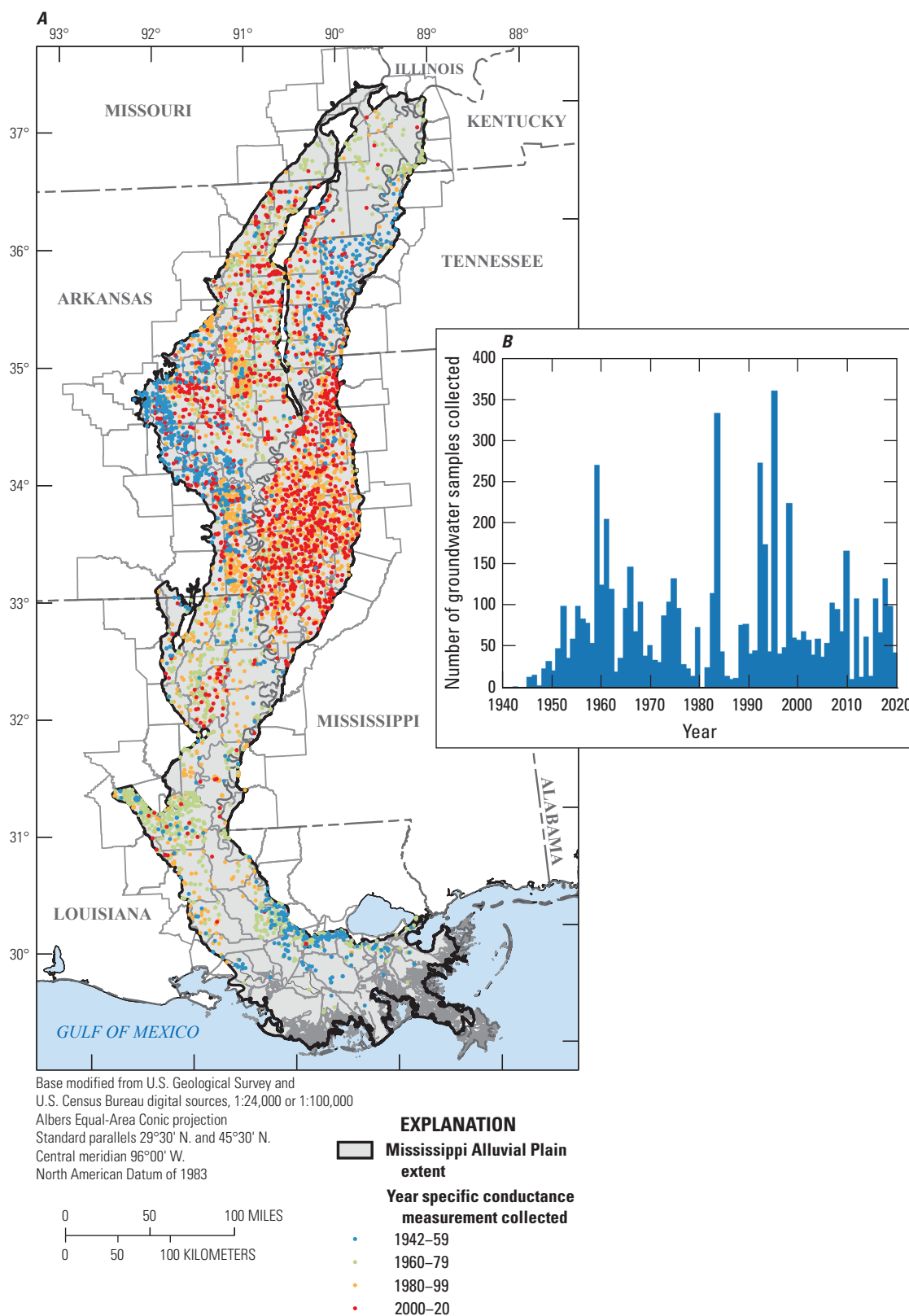


Figure 2. A, Spatiotemporal distribution of groundwater samples collected across the Mississippi Alluvial Plain region 1942–2020 and B, number of samples collected in each year used in the models to produce predictions of specific conductance for the region. Data from U.S. Geological Survey (2021) and Killian and Knierim (2023).

Study Area

The Mississippi Alluvial Plain region covers parts of seven States, extending as far north as Illinois and south to the Gulf of Mexico (fig. 1). The extent of the Mississippi Alluvial Plain covers about 44,200 square miles and can be divided into seven generalized subregions to allow for subregional comparisons of data (Ladd and Travers, 2019) (fig. 1A). The subregions are bounded by major streams or other ground-water divides, such as Crowleys Ridge. For this study, the southernmost subregion, the Deltaic and Chenier Plains, was expanded to include alluvial deposits along the Mississippi River between New Orleans and Lake Pontchartrain, La., and associated water-quality data (Killian and Knierim, 2023).

The alluvial aquifer is coextensive with the Mississippi Alluvial Plain region and overlies geologic units deposited by transgressive and regressive events during the Cretaceous and Tertiary Periods (Clark and Hart, 2009). The alluvial aquifer comprises Quaternary sands, silts, and clays deposited by the ancestral Mississippi and Ohio Rivers (Renken, 1998; Arthur, 2001). The thickness of the alluvial aquifer varies across the study area and is typically between 130 and 280 feet (ft) (Hart and others, 2008). In areas where the alluvial aquifer overlies subcropping aquifer units, the bottom of the alluvial aquifer is not well defined, and different sources provide different thickness estimates (Torak and Painter, 2019a; James and Minsley, 2021). A thin layer of shallow surficial confining material of silt, clay, and fine sand is present in the alluvial aquifer but may be absent in some areas (Renken, 1998). The surficial confining material is not uniform and is thought to impede recharge to the alluvial aquifer where present (Renken, 1998). The alluvial aquifer is above a large plunging syncline that dips toward the Gulf of Mexico. There are several subcropping aquifer hydrogeologic units that are Tertiary in age that have been mapped beneath the alluvial aquifer in the Mississippi embayment part of the study area (Hart and others, 2008).

Methods

Response Variable

Response variable data included spatially distributed measurements of SC collected from domestic wells and from public-supply, irrigation, and observation wells owned by several Federal, State, and local cooperators. SC water-quality data from 6,626 wells were obtained from the U.S. Geological Survey National Water Information System (U.S. Geological Survey, 2021). The Mississippi Department of Environmental Quality provided data from an additional 888 wells with SC water-quality measurements for a total of 7,514 wells (Killian and Knierim, 2023).

Wells that were less than 400 ft below the land surface were used in the response variable dataset, reducing the number of wells available for use in the models from 7,514 to

6,042 for the final dataset (Killian and Knierim, 2023) (fig. 3). Wells that were greater than 400 ft deep were not used because the depth of investigation of AEM survey data is generally 400 ft deep across the study area (James and Minsley, 2021). Wells completed in the alluvial aquifer and deeper Tertiary aquifers were included to allow the model freedom to predict without the constraint of aquifer designation because the base of the alluvial aquifer is not well defined in some areas. The median well depth was 114 ft below land surface (Killian and Knierim, 2023). A total of 987 wells did not have a defined bottom of well screen. Most of the wells were completed in the alluvial aquifer (number of samples [n] = 5,673), 98 wells had no aquifer designation, 49 wells were completed in terrace deposits, and the remaining wells were completed in deeper aquifers (n = 222).

Observed SC in the filtered dataset ranged from 4.7 to 23,400.0 microsiemens per centimeter at 25 degrees Celsius ($\mu\text{S}/\text{cm}$) (natural log 1.55 to 10.06) and was positively skewed (fig. 3) (Killian and Knierim, 2023). The median observation was 632 $\mu\text{S}/\text{cm}$ (natural log 6.45). BRT models split the response variable data into trees that minimize variance, which will allow higher values to have a greater effect on variance. Because of how BRT models handle response variable data and because the data are positively skewed, the data were log-transformed using the natural log to reduce the relative importance of the higher values and increase the predictive accuracy for the bulk of the response variable data.

The caret package in R (Kuhn and others, 2019) was used to randomly split the response variable data into two groups: a training dataset and a holdout dataset. The training data consisted of 4,835 measurements (about 80 percent) and were used to tune the model; the remaining 1,207 measurements (about 20 percent) were held out to evaluate model performance.

Groundwater wells across the Mississippi Alluvial Plain region have varying numbers of samples ranging from one sample to continuous monitoring throughout a year or across several decades. If more than one observation existed for a well, the most recent measurement was used in model tuning and prediction. To increase the number of SC measurements for the model, samples from 1942 to 2020 were used (fig. 2). Censored values, or measurements where the value is only partially known (for example, below the detection limit of an instrument), were not included in the response variable dataset (n = 7). Some subregions of the Mississippi Alluvial Plain region—such as the Atchafalaya and the Deltaic and Chenier Plains (fig. 1A)—have not been sampled for several decades, whereas some subregions—such as the Cache and the Delta (fig. 1A)—are lacking historical measurements (fig. 2A). The number of SC measurements collected across the Mississippi Alluvial Plain region changed throughout time (fig. 2B). Because only the most recent SC measurement at each well was used in both models, information about changes in SC through time was effectively removed, and this modeling effort assumed no change in water quality over time at a well.

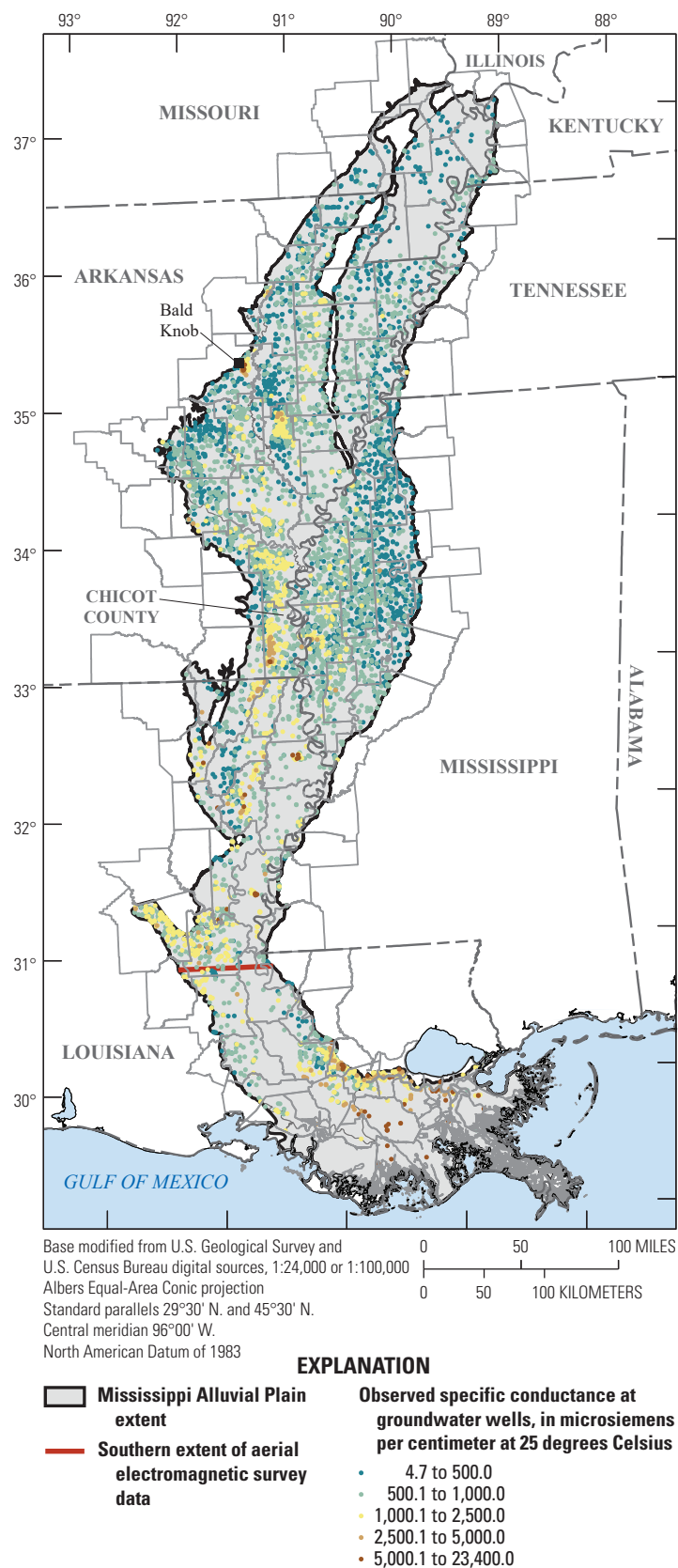


Figure 3. Specific conductance observed at groundwater wells screened in the Mississippi River Valley alluvial aquifer and deeper aquifers in the Mississippi Alluvial Plain region, 1942–2020 (sample size = 6,042), used to train two boosted regression tree machine-learning models. Data from U.S. Geological Survey (2021) and Killian and Knierim (2023).

Explanatory Variables

Two BRT models were developed to understand the influence of geophysical explanatory variables on SC predictions: one model without geophysical variables (the “baseline” model) and one model with geophysical variables from an aquifer-wide AEM survey (the “geophysical” model). SC predictions were compared between the BRT models to quantify the change in model performance and the sensitivity of the models to geophysical variables (Knierim and others, 2020; Stackelberg and others, 2021).

Both models initially included 29 explanatory variables describing well construction information; surficial characteristics such as soils and geomorphology; positional information such as latitude, longitude, generalized subregions, and multi-order hydrologic position; depth to the water table; and year of SC measurement (table 1, at back of report). The geophysical model included the same 29 explanatory variables as the baseline model and an additional 19 geophysical variables (for a total of 48). Geophysical explanatory variables included bulk electrical resistivity (average electrical resistivity through the depth of investigation), which is sensitive to groundwater salinity and aquifer texture. Additional variables interpreted from the AEM survey included surficial confining unit top and bottom altitudes, the estimated base of the alluvial aquifer, and aquifer connectivity with streams (table 1, at back of report). More information about each explanatory variable, including name, description, dataset group, and source information, can be found in table 1 (at back of report) and Killian and Knierim (2023).

Explanatory variables were attributed to well points by using Python (Python Software Foundation, 2021) to execute a point extraction from the coincident raster cell, calculate averages for a buffer around the well, or extract a value from the depth coincident with the well screen. Point and buffer extractions were completed with the zonepy package (Clark and others, 2019). Point extraction was used to maintain high resolution of the source data, where applicable (table 1, at back of report), and for categorical explanatory variables. To account for spatial variation around a well, continuous explanatory variables were attributed using a 500-meter (m) buffer around the well. Well buffers also serve as surrogates for the contributing area (Johnson and Belitz, 2009). Two of the geophysical variables—aquifer resistivity and facies class—were attributed using the midpoint depth of the well screen and the corresponding coincident raster cell and depth slice from the AEM survey data. If well depth was unknown, resistivity and facies class were not assigned.

The year of SC measurement collection was included in the BRT models as an explanatory variable because the SC dataset spanned 78 years. Ultimately, time is inherently included in the model because of the decades over which SC measurements were collected. By explicitly including the year of measurement collection as an explanatory variable, the effect of time on SC predictions could be determined.

Although the BRT models can make SC predictions for any specific year between 1942 and 2020, the models were not developed to predict SC through time. The explanatory variable dataset generally includes variables that capture conditions around 2018 (Killian and Knierim, 2023). To fully incorporate a time component in the BRT models and predict SC through time, explanatory variables that capture changes over time would also need to be included in model training.

Random noise was also included as an explanatory variable. The BRT models can determine the relative importance of each explanatory variable. Random noise was used as a benchmark to remove explanatory variables that were less influential than noise, which helped to improve model run time without impacting model performance. This method of including random noise for variable reduction is novel; however, similar approaches have been used (Ljunggren and Ishii, 2021). The random noise raster was generated with the ArcMap version 10.8 Create Random Raster tool available in the Spatial Analyst extension (Esri, 2020). The seed value was set to 2,640, and the Random Generator type was set to the default ACM599. Values in the random noise raster range from 0 to 1.

Geophysical Data Collection in the Mississippi Alluvial Plain Region and Incorporation in BRT Models

A unique aspect of the SC predictions from BRT models is incorporation of regional geophysical datasets, which provides information about the hydrogeology of the Mississippi Alluvial Plain region (Minsley and others, 2021). The geophysical data include an AEM survey, which measures resistivity signatures of water-bearing geologic materials (Smith and others, 2007). AEM survey data can help to map the thickness and extent of an aquifer. At the time of this study, the AEM survey data collected in the Mississippi Alluvial Plain region in 2019–21 covered 14,932 line-miles in the northern part of the study area and included shallow (less than about 330 ft) and deep (up to about 980 ft) estimates of inverted resistivity (mres) (James and Minsley, 2021; Minsley and others, 2021). Mres was further interpreted to estimate the bottom of the alluvial aquifer (mrvaBaseElv) and the presence or absence of surficial confining material (cnf_class). The AEM survey also included an CS-3 cesium vapor magnetometer (Scintrex Ltd., Concord, Ontario, Canada) to measure magnetic intensity (magint) and an RS-500 spectrometer radiochemical sensor (Radiation Solutions Inc., Mississauga, Ontario, Canada) to detect near-surface radioelements (predominantly potassium, uranium, and thorium) (Burton and others, 2021). These data provide three-dimensional information that has aided in the delineation of the base of the alluvial aquifer, locating where surficial confining material may be present or absent, and evaluating the degree of potential hydraulic connectivity between the alluvial aquifer

and deeper subcropping aquifers (James and Minsley, 2021). At the time of this study, AEM coverage extended from the northern extent of the Mississippi Alluvial Plain region to approximately 31 degrees north latitude (figs. 3, 4), meaning that the AEM survey data did not cover the full Mississippi Alluvial Plain region. The authors sought to test how a BRT model would perform with an incomplete explanatory variable dataset.

Model Development

BRT models were developed in the R language (R Core Team, 2021) by leveraging the *caret* and *gbm* packages (Kuhn and others, 2019; Ridgeway, 2019). BRT is an ensemble-tree method that builds many trees with branches based on if/then statements that group data into similar homogenous groups with respect to the response variable (Kuhn and Johnson, 2013). The BRT method provides models that are more stable with better predictive performance than a single decision tree (Elith and others, 2008). The number and shape of trees in BRT models are controlled by five hyperparameters, which determine the model fit and complexity. Hyperparameters include the number of trees (*nt*), interaction depth (*id*; how deeply the trees split), shrinkage (*sh*; the proportion of each tree used as the model grows), minimum number of observations per node, and bagging fraction (bootstrap aggregation) (Elith and others, 2008; Kuhn and Johnson, 2013; Ridgeway, 2019).

Cross-validation (CV) tuning was used to tune the baseline and geophysical models and find the combination of hyperparameters with the best performance, or lowest root-mean-square error (RMSE) (Kuhn and others, 2019). Bagging fraction was held at the recommended value of 0.5 (Friedman, 2001) and minimum number of observations per node at 10. A tuning grid was constructed with the following ranges for the three other hyperparameters: *nt* (500–10,000 × 500), *id* (6–12 × 2), and *sh* (0.002–0.012 × 0.002). The tuning grid included 480 hyperparameter combinations (Killian and Knierim, 2023). Tenfold CV tuning was used, where 10 percent of the training data are used as a testing dataset to measure RMSE, and the process is repeated 10 times (Kuhn and Johnson, 2013). CV tuning was completed using the *caret* package (Kuhn and others, 2019) on a computer with 86 cores and was completed in approximately 20 minutes.

The BRT model with the lowest RMSE was considered the “best” model; however, the best model may be complex based on the hyperparameters, higher *nt*, *id*, and *sh*. Conversely, simpler models have lower *nt*, *id*, and *sh*. Models with high complexity may overfit to the training data and perform poorly when making new predictions. To find BRT models with similar predictive performance, models of lower complexity with RMSEs within one standard error (1SE) of the best model’s RMSE (Nolan and others, 2015) were inspected. Holdout data were used to evaluate 1SE model performance. Of the 480 hyperparameter combinations, there

were 127 1SE baseline models and 172 1SE geophysical models. Because any 1SE model may provide reasonable predictions and could be selected as the “final” model, a “favorite” model was selected somewhat subjectively from the pool of 1SE models—one each for the baseline and geophysical models—favoring a simpler model (relatively low complexity controlled by the hyperparameters) and lower RMSE. Once the favorite model was selected, explanatory variables less important than the noise explanatory variable were removed on the basis of the assumption that random noise is insignificant to the model, and the model was retrained to obtain the “final” model. The final model was used to make SC predictions at wells and across the aquifer for both the baseline and geophysical models.

Ensemble-tree models, such as BRT, tend to underpredict high observations and overpredict low observations (Zhang and Lu, 2012; Belitz and Stackelberg, 2021). Empirical distribution matching (EDM) was used to correct for this bias and backtransform predictions from the natural log to original units (microsiemens per centimeter) (Belitz and Stackelberg, 2021). The EDM method calculates two empirical cumulative distribution functions for the training dataset: one for observed values and one for the ML predictions. Ordered pairs are matched between the two functions, which allow for linear interpolation between the points (Belitz and Stackelberg, 2021). As a result, if the training dataset includes the minimum and maximum observed values, then the range of EDM-corrected values will match the range of observations.

BRT models were inspected to better understand the importance and influence of explanatory variables for predicting SC. Variable importance was quantified using the *gbm* package in R (Ridgeway, 2019). Partial dependence plots (PDPs) quantify how a specific explanatory variable influences the response variable and were created using the *pdp* package in R (Greenwell, 2017). A PDP is generated from an average of individual condition expectation plots (Goldstein and others, 2015; Greenwell, 2017). These plots show the predicted value of the response variable (natural log of SC) on the y-axis if the explanatory variable of interest on the x-axis is changed over a range of observations and all other explanatory variables are kept at their observed or measured value (Elith and others, 2008; Greenwell, 2017; Knierim and others, 2020).

Prediction Mapping

SC predictions were made across the Mississippi Alluvial Plain region at multiple depths by using the final tuned BRT baseline and geophysical models and the “predict” function of the *gbm* R package (Ridgeway, 2019; R Core Team, 2021). Explanatory variable rasters were resampled to the 1-kilometer National Hydrogeologic Grid (Clark and others, 2018), and a flat file of the values (“raster stack”) was created. Predictions were made on the National Hydrogeologic Grid at 20-, 75-, 100-, 200-, and 300-ft depths below land surface. Prediction depths coincide with approximate depths to water at cones of

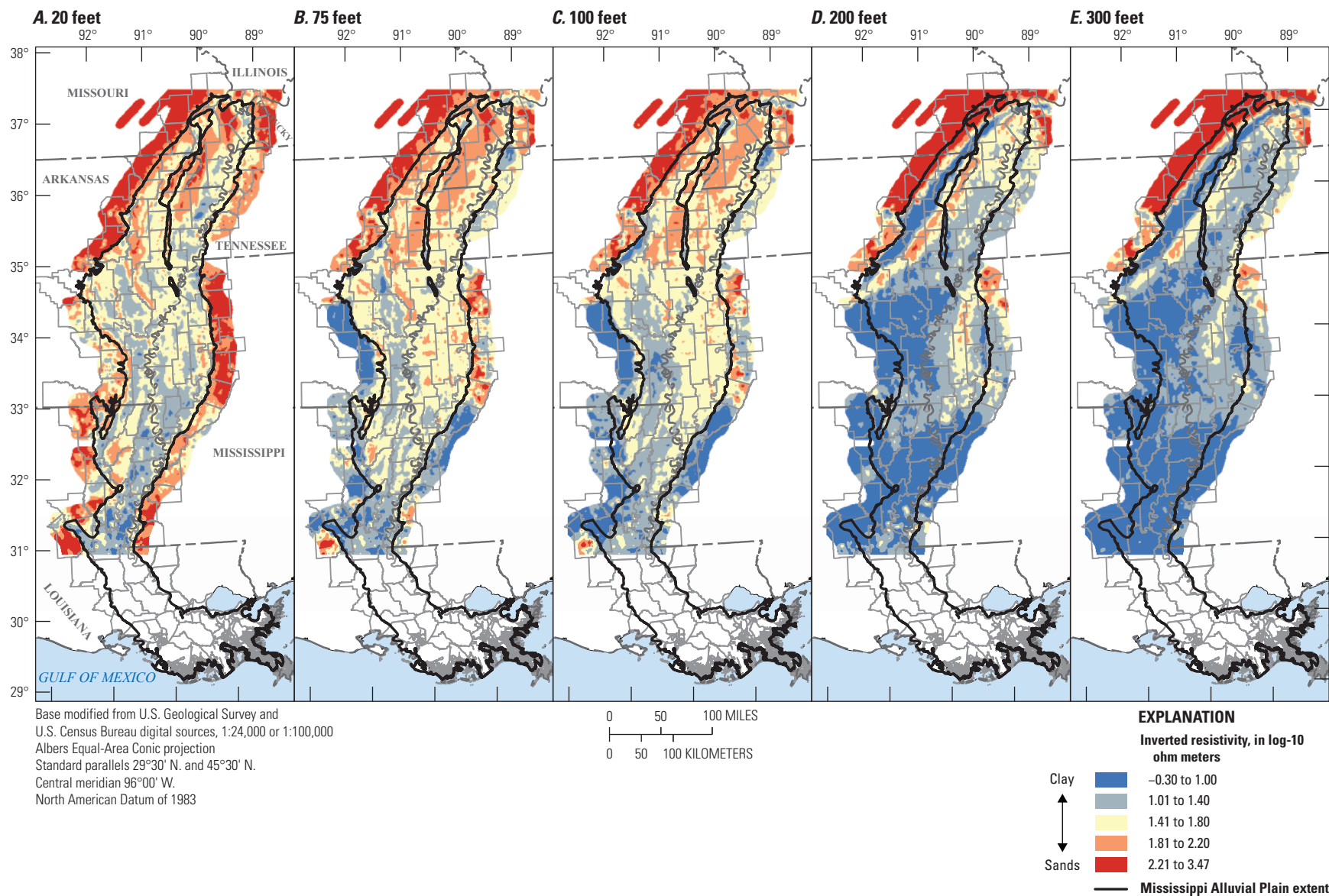


Figure 4. Inverted resistivity estimated from aerial electromagnetic survey data collected during 2019–21 in the Mississippi Alluvial Plain region by depth slices of *A*, 20 feet, *B*, 75 feet, *C*, 100 feet, *D*, 200 feet, and *E*, 300 feet (survey data from James and Minsley [2021] and Minsley and others [2021]).

depression and the depth of investigation for the geophysical data (Minsley and others, 2021; McGuire and others, 2023). Continuous prediction surfaces were generated for the entire Mississippi Alluvial Plain extent, and a water-level mask was applied to remove predictions that were made above the 2018 alluvial aquifer potentiometric surface (McGuire and others, 2020a). The year 2018 was chosen for the prediction year because it was coincident with the geophysical data collection (James and Minsley, 2021; Minsley and others, 2021) and the spring 2018 potentiometric surface (McGuire and others, 2020b) for the alluvial aquifer (spring is when water levels are assumed to be at maximum recovery, before the start of the main irrigation season). Predictions were bias corrected with the EDM method and backtransformed to units of microsiemens per centimeter.

Uncertainty

The 1SE models were used to quantify the range of SC predictions at the 200-ft prediction depth. Any of the 1SE models may provide a reasonable prediction; leveraging all of the 1SE model predictions provides a way to quantify the variability in SC predictions across the Mississippi Alluvial Plain region. This variability is used to express model uncertainty, though confidence intervals were not calculated. Prediction maps were generated for the 127 1SE baseline models and 172 1SE geophysical models. The rasters were then used to calculate the 5th percentile, 50th percentile (median), 95th percentile, and difference between the 95th and 5th percentiles of SC predictions for both models at each cell.

SC predictions at the 5th percentile, median, and 95th percentile were used to provide qualitative classifications of uncertainty relative to a threshold. TDS are used to define brackish and saline groundwaters, and SC and TDS show a highly correlated linear relation across the Mississippi Alluvial Plain region (Knierim and others, 2020). Therefore, TDS were calculated from SC predictions by using the correlation workflow in Knierim and others (2020). The TDS prediction rasters were then used to define where the models would very likely or likely predict TDS greater than ($>$) or less than ($<$) 1,000 mg/L, the brackish water threshold. The following logic was used:

If TDS-5th $<$ 1,000 mg/L and TDS-median $<$ 1,000 mg/L and TDS-95th $<$ 1,000 mg/L, then TDS very likely $<$ 1,000 mg/L.

If TDS-5th $<$ 1,000 mg/L and TDS-median $<$ 1,000 mg/L and TDS-95th $>$ 1,000 mg/L, then TDS likely $<$ 1,000 mg/L.

If TDS-5th $<$ 1,000 mg/L and TDS-median $>$ 1,000 mg/L and TDS-95th $>$ 1,000 mg/L, then TDS likely $>$ 1,000 mg/L.

If TDS-5th $>$ 1,000 mg/L and TDS-median $>$ 1,000 mg/L and TDS-95th $>$ 1,000 mg/L, then TDS very likely $>$ 1,000 mg/L.

Results

BRT Model Results and Performance

The “final” BRT models underpredicted high observations and overpredicted low observations, which is typical of decision tree ML models (Zhang and Lu, 2012) (fig. 5A, B). When predictions were bias corrected with the EDM method, the predicted SC values matched the range of observed values because the training dataset included the minimum and maximum observations (fig. 5C, D). For the baseline model, EDM bias-corrected SC predicted at each well ranged from 4.7 to 23,400.0 μ S/cm (natural log 1.55 to 10.06), with a median prediction of 634.0 μ S/cm (natural log 6.45) (Killian and Knierim, 2023). For the geophysical model, the EDM bias-corrected SC predicted at each well ranged from 4.7 to 23,400.0 μ S/cm (natural log 1.55 to 10.06), with a median prediction of 631.5 μ S/cm (natural log 6.45) (Killian and Knierim, 2023). The EDM-bias correction method (Belitz and Stackelberg, 2021) provided better matches between the observed and predicted values for both models.

Overall BRT model performance, reported as coefficient of determination (R^2) and RMSE, was similar for the baseline and geophysical models (fig. 5, table 2). The EDM bias-corrected baseline model training R^2 and holdout R^2 values were 0.896 and 0.617, and the EDM bias-corrected geophysical model training R^2 and holdout R^2 were 0.906 and 0.626 (fig. 5) (Killian and Knierim, 2023). There was a slight increase in model accuracy based on R^2 value of holdout data for the geophysical model (fig. 5). Training data predictions generally had higher accuracy (lower RMSE) than did holdout data predictions, which is typical for BRT models. Training data accuracy is typically higher than holdout data accuracy because training data were used to develop the model and holdout data were used to assess how well the model can predict using new data. Holdout RMSE values ranged from 0.42 to 0.45 (table 2) (Killian and Knierim, 2023).

Each Mississippi Alluvial Plain subregion had a different number of SC measurements, and the range and accuracy of SC predictions varied by subregion (fig. 6). Training R^2 values were similar among subregions, ranging from 0.79 to 0.92 (Killian and Knierim, 2023). Holdout R^2 values ranged from 0.26 to 0.88 (Killian and Knierim, 2023). The geophysical model had a higher R^2 value than the baseline model for the holdout data for all subregions except the Atchafalaya and Grand Prairie. Holdout R^2 values for both models were closest in the Delta subregion, where the models had the highest number of SC measurements. Holdout R^2 values were highest for both models in the Boeuf subregion. The Saint Francis subregion had the lowest holdout R^2 value for the baseline model, where SC is generally lower relative to the rest of the dataset. The Atchafalaya subregion had the lowest holdout R^2 value for the geophysical model. Holdout R^2 values differed the most in the Deltaic and Chenier Plains subregion, where SC is highest and the number of observations is lowest.

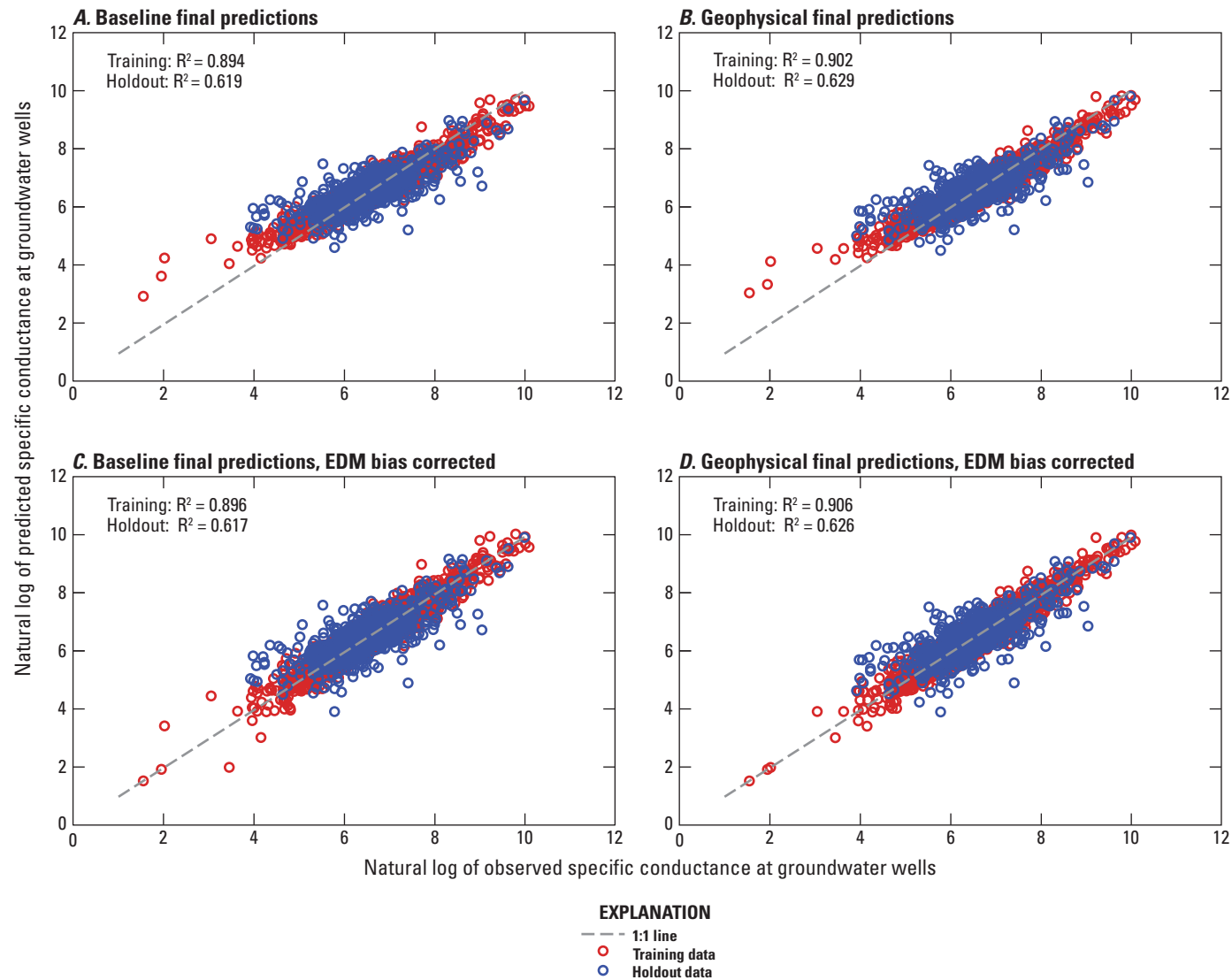


Figure 5. Comparison of observed and predicted specific conductance values at groundwater wells by training and holdout datasets with coefficient of determination (R^2) for the final *A*, baseline and *B*, geophysical models and for the final *C*, baseline and *D*, geophysical models bias corrected with the empirical distribution matching (EDM) method for specific conductance data collected 1942–2020 in the Mississippi Alluvial Plain region. Data from U.S. Geological Survey (2021) and Killian and Knierim (2023).

Table 2. Hyperparameters for boosted regression tree (BRT) model type for the empirical distribution matching (EDM) bias-corrected, final (variable-reduced), favorite, and best training models for predicting specific conductance at groundwater wells in the Mississippi Alluvial Plain region.

[Datasets from U.S. Geological Survey (2021) and Killian and Knierim (2023). n, number of samples; no. explanatory vars, number of explanatory variables; id, interaction depth; mo, minimum number of observations per node; sh, shrinkage; nt, number of trees; RMSE, root-mean-square error]

BRT model type	n, training dataset	n, holdout dataset	Model	no. explanatory vars	Hyperparameter				Model complexity	Holdout RMSE
					id	mo	sh	nt		
Baseline	4,835	1,207	EDM	22	6	10	0.012	9,500	8	0.45
			Final	22	6	10	0.012	9,500	8	0.44
			Favorite	29	6	10	0.012	9,500	8	0.44
			Best	29	12	10	0.006	10,000	118	0.43
Geophysical	4,835	1,207	EDM	33	6	10	0.01	10,000	7	0.44
			Final	33	6	10	0.01	10,000	7	0.43
			Favorite	48	6	10	0.01	10,000	7	0.43
			Best	48	12	10	0.01	8,000	153	0.42

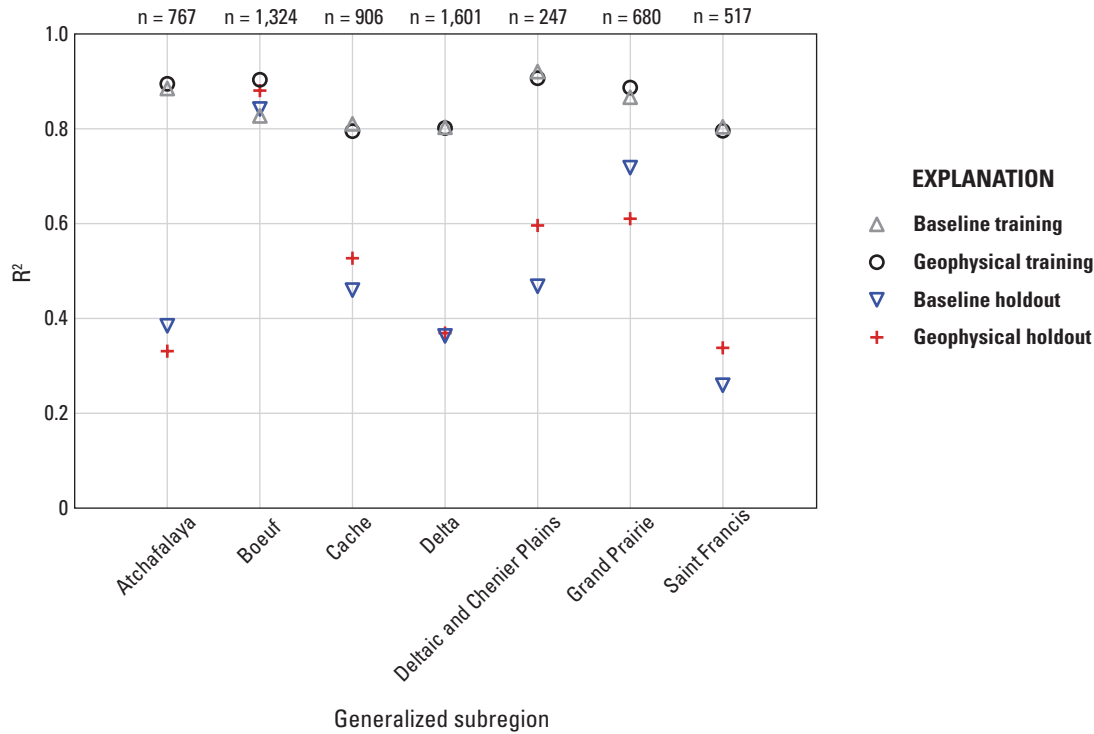


Figure 6. Baseline and geophysical model training and holdout coefficients of determination (R^2) values for predictions of specific conductance at groundwater wells by generalized subregion of the Mississippi Alluvial Plain region. Model datasets (1942–2020) from U.S. Geological Survey (2021) and Killian and Knierim (2023); n, number of samples.

When the models are compared by depth slice, SC predictions tend to be higher and more accurate for the geophysical model compared to the baseline model. Holdout R^2 values generally increased with depth for both models (fig. 7). Both models also had holdout R^2 values less than 0.2 for wells in the 20-ft depth slice (0–20 ft below land surface). Wells less than 20 ft deep (that is, closest to the water table) were the least well defined ($n = 78$ wells). The geophysical holdout had a higher R^2 value compared to the baseline holdout for all depths except for wells in the 75-ft depth slice.

The geophysical model tended to predict higher maximum SC values at each depth slice compared to the baseline model, ranging from 30 to 200 $\mu\text{S}/\text{cm}$ higher, although the median predictions were nearly identical (fig. 8A). Once the models were bias corrected using the EDM method, the difference in maximum predictions among depth slices between the baseline and geophysical models was less noticeable (fig. 8B). Therefore, the geophysical explanatory variables may provide additional information to the BRT model to predict high SC values prior to bias correction.

SC Raster Predictions and Uncertainty

Both the baseline and geophysical models predicted a general increase in SC from north to south, with the highest SC around the Gulf of Mexico (pl. 1). Additionally, both models predicted high SC in areas of known high salinity: around Iberville Parish, La. (Welch and Hanor, 2011); in southeastern Arkansas (Larsen and others, 2021); and around White County, Ark. (Bryant and others, 1985) (figs. 1, 3). The

baseline model predicted higher SC near the Gulf of Mexico compared to the geophysical model; frequency of predictions increased between 5,000.1 and 28,000.0 $\mu\text{S}/\text{cm}$ in the baseline model (pl. 1, A). The geophysical explanatory variables do not extend to the Gulf of Mexico (figs. 3, 4), so a true one-to-one evaluation of model predictions cannot be quantified. However, a one-to-one comparison of models with identical explanatory variable extents coincident with AEM survey data coverage was explored, and no significant impacts to model performance were discovered because of the nature of how BRT models function. The comparison between models with differing explanatory variable extents was sufficient to assess the impact of AEM survey data on the model, as well as to assess the ability of a BRT model to predict where explanatory variable data are lacking. The extent of AEM survey data coverage created a horizontal linear feature in the geophysical model for all depth predictions (pl. 1, B). The difference in predictions between the baseline and geophysical models increased at the 200- and 300-ft depth slices compared to shallower predictions (pl. 1, C).

The geophysical model predicted higher SC with increasing depth compared to the baseline model, especially above 31 degrees north latitude, where the geophysical datasets were collected (figs. 8, 9, pl. 1). For the 20-, 75-, and 100-ft depth slices, the baseline model predicted a larger interquartile range and higher median than did the geophysical model (fig. 9). The interquartile ranges for the 200- and 300-ft depth slices for both models were similar, but the geophysical model had higher medians and ranges of predictions (fig. 9). This difference between the baseline and geophysical model predictions is more noticeable in areas of known high SC where

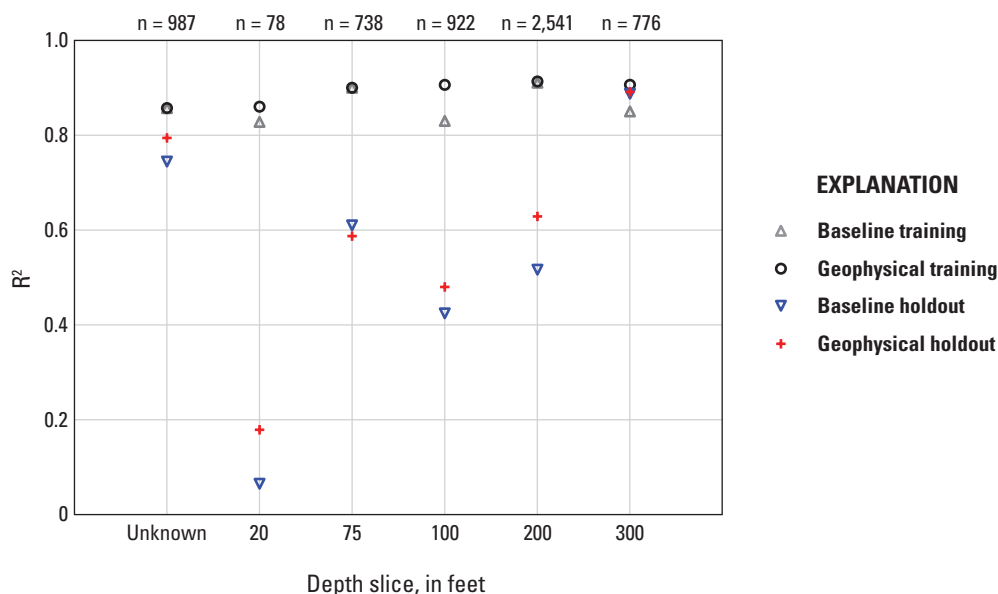


Figure 7. Baseline and geophysical model training and holdout coefficients of determination (R^2) values for predictions of specific conductance by depth slice at groundwater wells in the Mississippi Alluvial Plain region. Model datasets (1942–2020) from U.S. Geological Survey (2021) and Killian and Knierim (2023); n, number of samples.

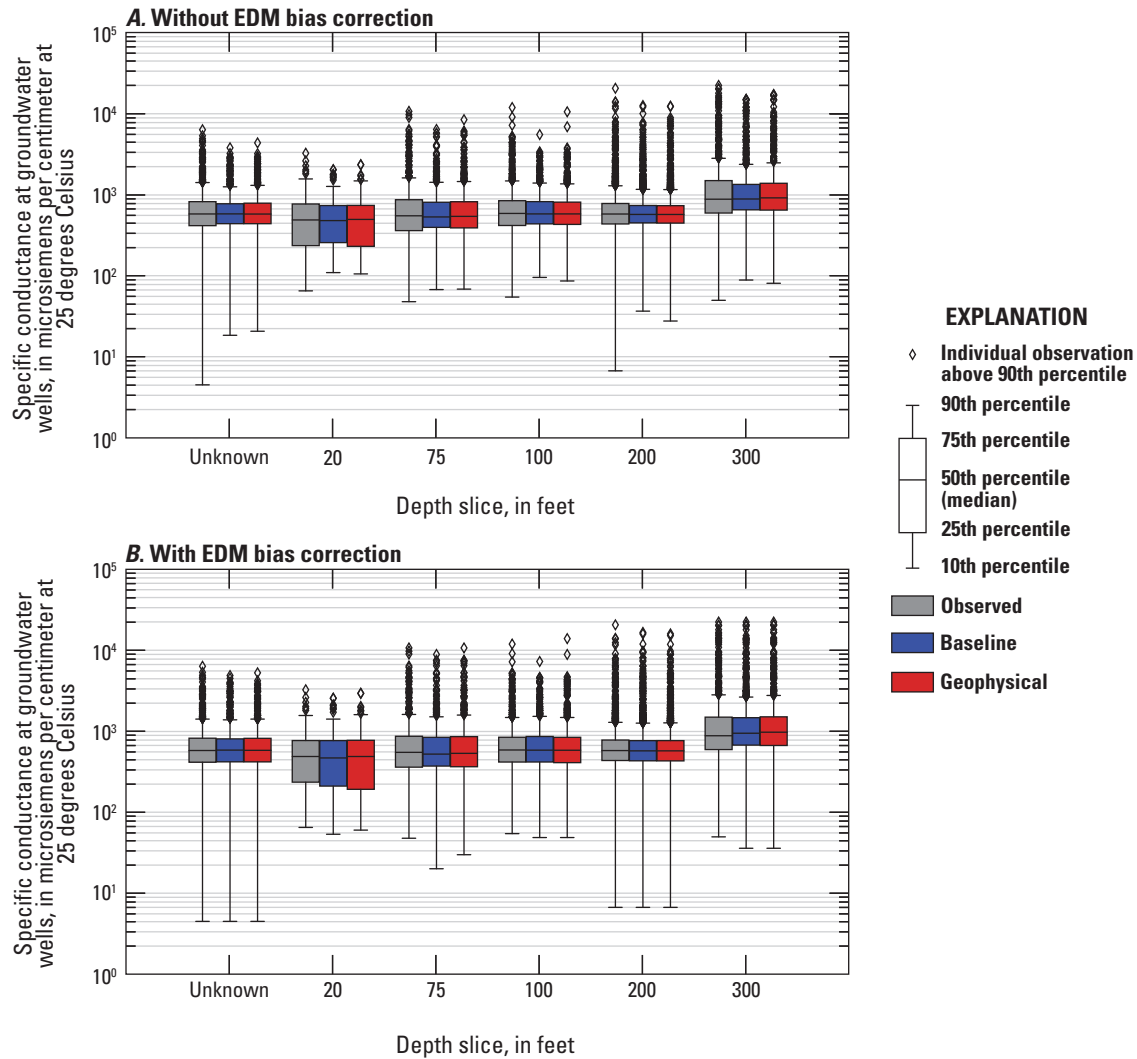


Figure 8. Observed specific conductance 1942–2020 at groundwater wells in the Mississippi Alluvial Plain region (U.S. Geological Survey, 2021; Killian and Knierim, 2023) by depth slice with predictions of specific conductance from the baseline and geophysical models *A*, without and *B*, with bias correction using the empirical distribution matching (EDM) method (boxplots created with Python seaborn package [Waskom and others, 2021]).

geophysical explanatory variables are complete (that is, no missing data), such as southeastern Arkansas and northeastern Louisiana (pl. 1).

Uncertainty in SC predictions was high where observed SC was high and where there was a greater range in SC predictions among all ISE models in both the baseline and geophysical models (fig. 10*A, B*). High relative uncertainty where the response variable is high has been noted in other ML models (Knierim and others, 2020). The importance of quantifying uncertainty in model predictions depends on how the predictions will be used; for example, whether the modeling effort is focused on accuracy for low or high observed values. High groundwater salinity can limit drinking water and irrigation uses; therefore, qualitative uncertainty was mapped to show where the model is very likely or likely to predict greater or less than the brackish groundwater threshold of 1,000 ppm TDS (fig. 10*C, D*) (U.S. Geological Survey, 2013). Overall, there were few areas where either model does not predict fresh versus brackish groundwater with a high level of confidence. Confidence was determined by whether the 5th percentile (considered the lowest prediction), median, and 95th percentile (considered the highest prediction) all predicted SC greater or less than 1,000 ppm TDS.

Explanatory Variable Influence

Both the baseline and geophysical models had fewer explanatory variables in the final model compared to the potential explanatory variables used in initial model tuning

(table 2). Twenty-two of the 29 explanatory variables were more important than noise and retained in the final baseline model (fig. 11*A*). The geophysical data added 19 potential explanatory variables. Thirty-three of the 48 explanatory variables, including 12 explanatory variables from the AEM survey, were more important than noise and retained in the final geophysical model (fig. 11*B*). The same six explanatory variables were found to be less influential than noise in both models, including variables that describe soil texture and the age of surficial sediments (fig. 11).

PDPs for continuous explanatory variables shared between both models show a similar pattern for predicting SC (fig. 12). For example, SC was predicted to be lowest at approximately -92 degrees longitude and highest at approximately -91 degrees longitude in both models (fig. 12*E*), despite longitude being the first and third influential for the baseline and geophysical models, respectively (fig. 11). The baseline model PDPs generally predicted a higher range of SC than did the geophysical model PDPs for the shared explanatory variables (fig. 12). The depth of the well screen (BOT) was one explanatory variable that showed a different relation between the two models; SC increased with depth in the baseline model and decreased with depth in the geophysical model (fig. 12*B*). Eleven continuous explanatory variables (fig. 12*U–AE*) were unique to the geophysical model, and the variable with the largest range of SC predictions was mres (fig. 12*Z*).

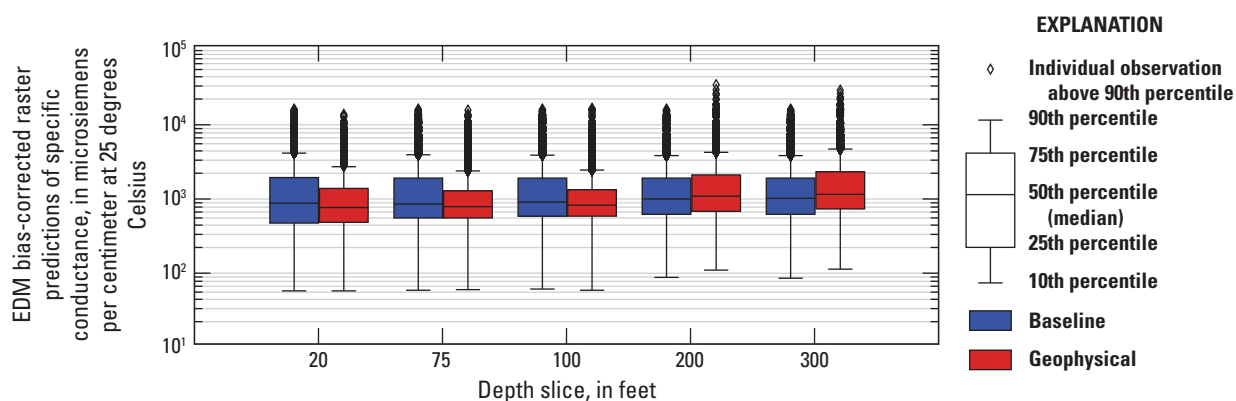


Figure 9. Raster predictions of specific conductance at groundwater wells in the Mississippi Alluvial Plain region by depth slice for the baseline and geophysical models (boxplots created with Python seaborn package [Waskom and others, 2021]). Model datasets (1942–2020) from U.S. Geological Survey (2021) and Killian and Knierim (2023); EDM, empirical distribution matching.

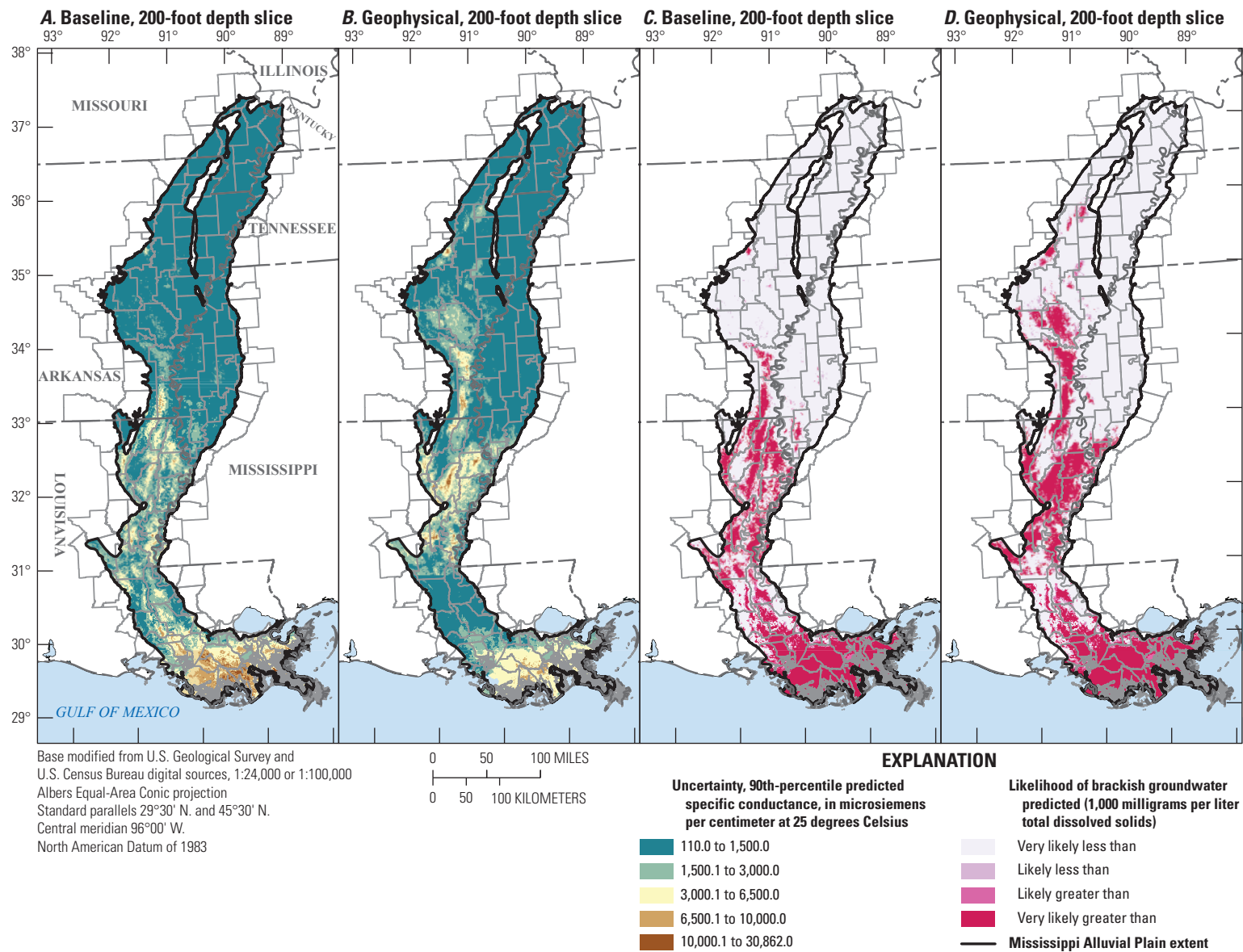


Figure 10. Quantification of specific conductance prediction uncertainty at the 200-foot depth slice for the *A*, baseline and *B*, geophysical models and likelihood of brackish groundwater (U.S. Geological Survey, 2013) predicted at the 200-foot depth slice by the *C*, baseline and *D*, geophysical models for the Mississippi Alluvial Plain region.

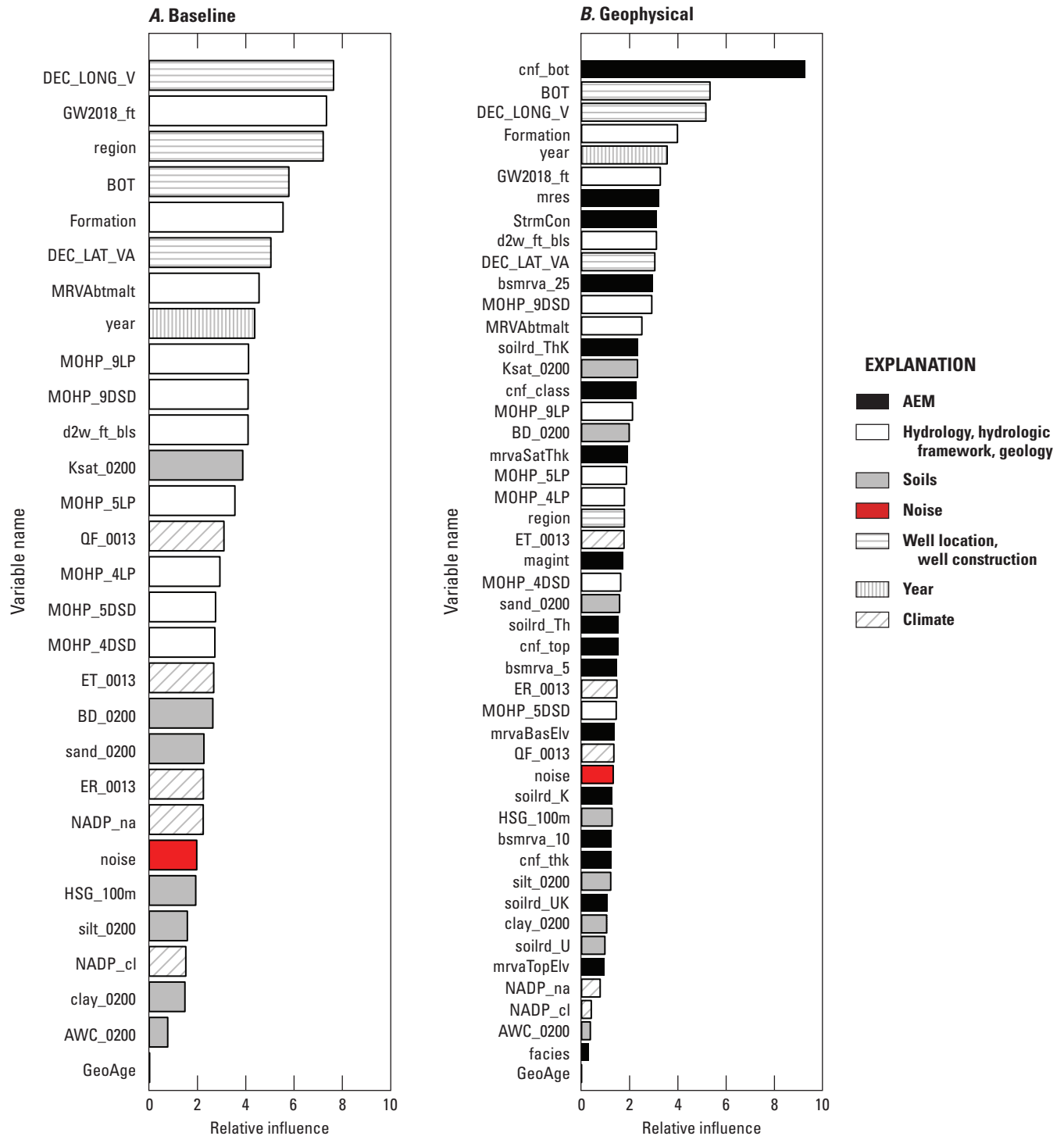


Figure 11. Relative influence of explanatory variables in the *A*, baseline model (excluding aerial electromagnetic [AEM] explanatory variables) and *B*, geophysical model (including AEM explanatory variables) for predicting groundwater specific conductance at wells in the Mississippi Alluvial Plain region. The noise variable and all explanatory variables below noise were removed from the final models. See [table 1](#) (at back of report) for descriptions and units of variables.

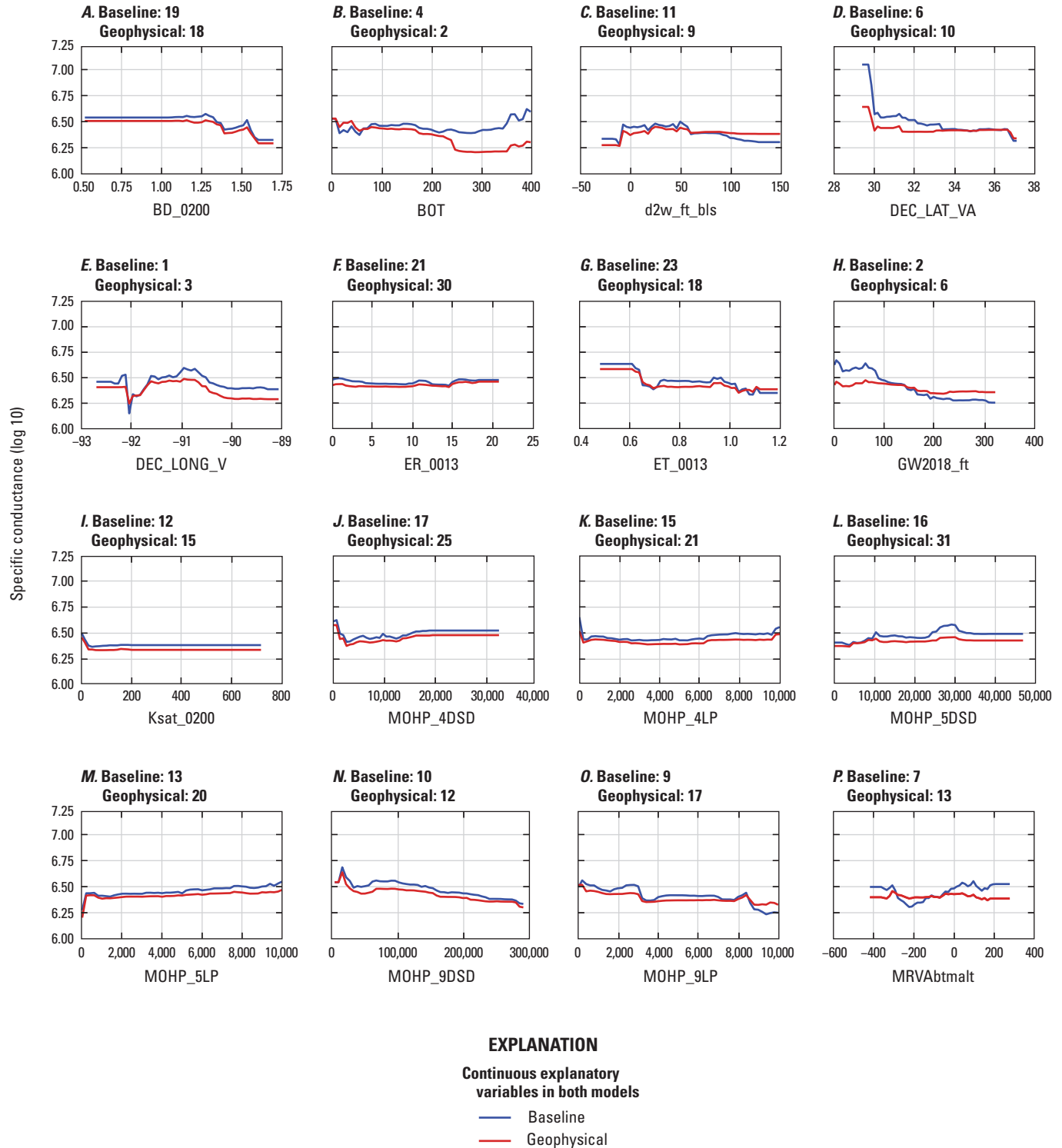


Figure 12. Partial dependence plots of continuous explanatory variables used in the final baseline and geophysical models for predicting groundwater specific conductance at wells in the Mississippi Alluvial Plain region. Numbers in individual graph titles indicate the relative influence ranking of the variable (fig. 11). See table 1 (at back of report) for explanatory variable definitions and units.

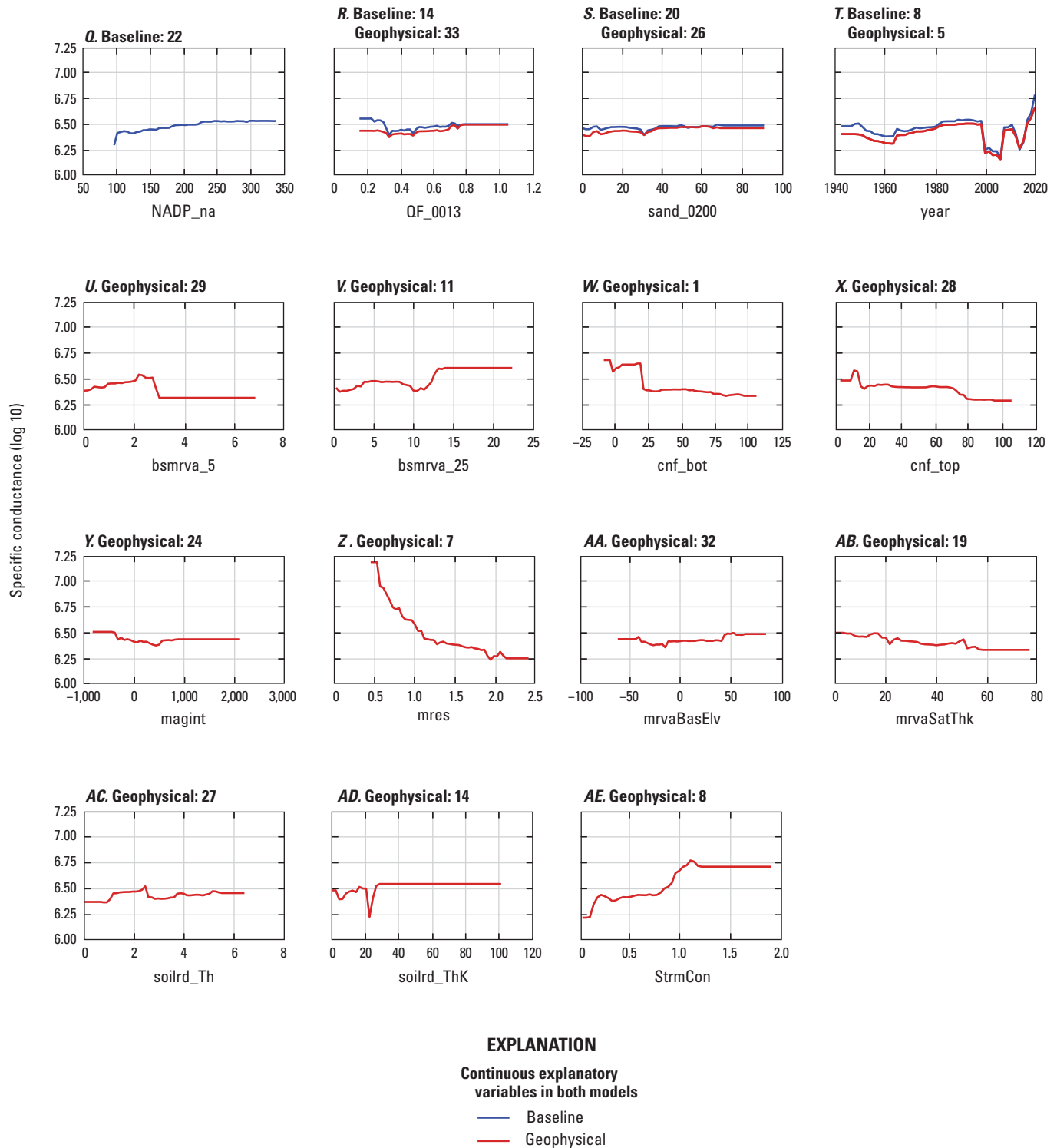


Figure 12. Partial dependence plots of continuous explanatory variables used in the final baseline and geophysical models for predicting groundwater specific conductance at wells in the Mississippi Alluvial Plain region. Numbers in individual graph titles indicate the relative influence ranking of the variable (fig. 11). See table 1 (at back of report) for explanatory variable definitions and units.— Continued

Discussion

This section discusses the differences and similarities between the predictions of the baseline and geophysical models. Additionally, the model results are discussed in the context of interpretations of the sources of salinity within the Mississippi Alluvial Plain region.

Missing Data Versus Structured Nulls in BRT Models

BRT models can make predictions even where explanatory variables are missing data because the decision trees are built with the missing data, or null values. Both the baseline and geophysical models were able to predict SC in the southern part of the Mississippi Alluvial Plain region, where there were fewer SC observations (fig. 3, pl. 1). In addition to the paucity of SC observations, geophysical data were not collected below approximately 31 degrees north latitude in the Mississippi Alluvial Plain region prior to or during model development (figs. 3, 4); therefore, there is a large area of non-randomly distributed (or “structured”) nulls. Because overall model performance was not impacted by the presence of a structured null, as assessed by a one-to-one model comparison with identical extents not described herein, the ability of a BRT model to predict with a structured null as a result of differing explanatory variable extents was tested. For wells in this area, the geophysical model had fewer explanatory variables on which to train. The large structured null created a linear feature in SC predictions for the geophysical model at approximately 31 degrees north latitude (pl. 1, B). Although BRT models can make predictions where there are randomly distributed missing data in explanatory variables, the BRT models do not appear to compensate for structured nulls.

SC was predicted to be lower in the geophysical model relative to the baseline model in the southern extent of the study area (pl. 1), despite the general trend of increasing observed SC toward the Gulf of Mexico (fig. 3). SC predictions north of the linear horizontal feature created by the structured null for the geophysical model were similar between the two models up to approximately the 200-ft depth slice (pl. 1). It cannot be determined which model would predict higher SC in the southern part of the Mississippi Alluvial Plain region if geophysical data were available for the full extent. Although predictive performance is similar between the two models in the Deltaic and Chenier Plains subregion (fig. 6), SC predictions from the geophysical model are likely not reasonable because of the large structured null coincident with missing geophysical explanatory variables.

Position in the Groundwater System Influences SC

The most influential explanatory variables for both the baseline and geophysical models indicate that position in the groundwater system is an important predictor of SC across the Mississippi Alluvial Plain region (fig. 11). A similar result was observed in a previous study in the Mississippi Alluvial Plain region where surrogates for position along groundwater flow paths were important predictors of SC and chloride (Knierim and others, 2020).

The top three most influential explanatory variables for the baseline model (DEC_LONG_V, GW2018_ft, and region [subregion]) are indicators for position in the groundwater system (fig. 11A). SC was predicted to be highest at approximately −91 degrees longitude (fig. 13A), which is the approximate center of the Mississippi Alluvial Plain and where the alluvial aquifer tends to be thickest (Clark and others, 2011). The 2018 alluvial aquifer potentiometric surface (fig. 13C) also reflects position in the system because groundwater altitude is higher in the northern part of the Mississippi Alluvial Plain region, where land surface altitudes are higher compared to sea level. SC was predicted to be high where the potentiometric surface is low (fig. 13C), which occurs toward the Gulf of Mexico. A similar north-to-south trend of increasing SC toward the Gulf of Mexico is evident in SC observations and mapped predictions (fig. 3, pl. 1). SC was also influenced by the subregion where the SC measurement was collected (figs. 11A, 13). The southernmost subregion, the Deltaic and Chenier Plains, tends to have higher SC because of saltwater influence from the Gulf of Mexico (Welch and Hanor, 2011).

The geophysical model was also influenced by explanatory variables that describe position in the groundwater system. For the geophysical model, the most influential explanatory variable was the bottom altitude of the surficial confining material (cnf_bot) (fig. 11B). Like groundwater altitude, the altitude of the bottom of surficial confining material tends to mimic land surface altitude and decreases from north to south. SC was predicted to be higher where the cnf_bot altitude was low, between 0 and 20 m (fig. 13B), which occurs in the southern half of the Boeuf and Delta subregions. The second most influential explanatory variable for the geophysical model was the bottom of the well screen relative to land surface (BOT) (fig. 13D). Generally, SC would be expected to increase along groundwater flow paths because dissolved solids increase with greater rock-water interaction; depth can serve as a surrogate for flow path (Moore and others, 2019). However, the geophysical model predicted lower SC with increasing well bottom screen depth, which is opposite of the baseline model (fig. 12B) (see below for discussion).

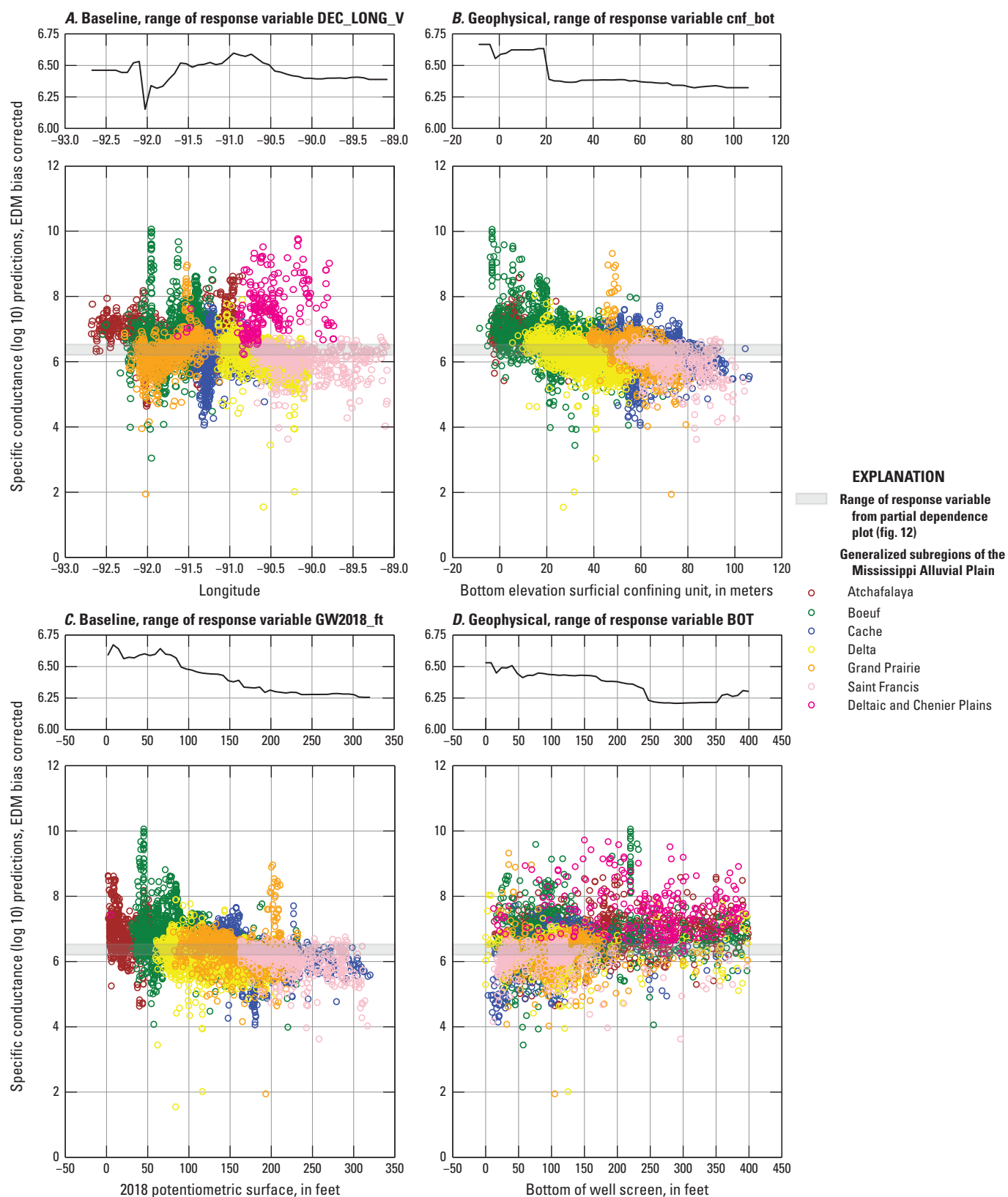


Figure 13. Comparison of the most influential explanatory variables in the *A*, baseline and *B*, geophysical models and the second most influential explanatory variables in the *C*, baseline and *D*, geophysical models and specific conductance predictions by subregion at groundwater wells in the Mississippi Alluvial Plain region. See [table 1](#) for explanatory variable definitions and units; EDM, empirical distribution matching.

Variation in Model Performance and SC Predictions With Depth

The baseline and geophysical BRT models had similar overall predictive performance despite being tuned with different explanatory variables (tables 1, 2, figs. 5, 7, 11). The predictive performance of both models was also similar to a previous study that predicted SC by using a BRT model in a smaller part of the study area (Knierim and others, 2020). The poor model performance at the 20-ft depth slice ($n = 78$) may be explained by the lack of SC observations near the water table and thus a lack of training data (fig. 7). Because of the poor model performance at this depth slice, the 20-ft depth prediction is likely inaccurate. More SC observations near the water table would be required to increase prediction accuracy at shallow depths; with the current training data, the BRT models should not be used to make predictions less than 75 ft deep.

Overall, the baseline model predicted similar SC at wells for all depth slices and across all raster depth slices, whereas the geophysical model predicted increasing median SC with depth at wells and raster depth slices (figs. 8, 9, pl. 1). Observed SC was not different among most depth slices, except for higher SC at the 300-ft depth slice (fig. 8), which likely includes wells below the alluvial aquifer. The baseline model included explanatory variables that predominantly describe surficial conditions across the Mississippi Alluvial Plain region (such as multi-order hydrologic position or soils) or two-dimensional position in the system (fig. 11A). BOT was the only three-dimensional explanatory variable in the baseline model and was an important predictor in both models (fig. 11). The geophysical model incorporated additional three-dimensional data including mres measurements that vary with depth within the system (figs. 4, 11B). With the addition of geophysical data, surficial variables decreased in importance in the geophysical model, but position in the groundwater system was still the predominant driving predictor of SC in the Mississippi Alluvial Plain region (fig. 11B). The addition of three-dimensional data in the geophysical model may explain why the median SC predictions at the 200- and 300-ft depth slices are higher than in the baseline model (fig. 9, pl. 1).

The PDPs for BOT show contrasting behavior between models in comparison to mapped predictions; SC was predicted to increase with well screen bottom depth in the baseline model, but it decreases in the geophysical model (fig. 12B). This contrasting relation may be related to the structured nulls in the geophysical model, which led to generally lower SC predictions in the geophysical model compared to the baseline model south of the geophysical data extent (pl. 1). Each PDP provides SC predictions if all other explanatory variables are held at their measured or observed values, except for BOT. The mapped predictions integrate spatial changes in all explanatory variables. Therefore, the BOT explanatory variable may modify the influence of other explanatory variables in predicting SC differently between the models. If AEM geophysical data were available for the

entire Mississippi Alluvial Plain extent, the BOT PDP for the geophysical model may display a trend similar to that for the baseline model.

Geophysical Model Explanatory Variables and Sources of Salinity

Areas of elevated SC in the Mississippi Alluvial Plain region have been suggested to be controlled by several possible sources of salinity, including ocean water; salt domes; connection with deeper, more saline subcropping aquifers; or surficial concentration through evaporation in areas of low infiltration controlled by the presence or absence of surficial confining units (Pettijohn and others, 1988; Williamson and others, 1990; Kresse and Clark, 2008; Kingsbury and others, 2014). The north-to-south trend of increasing SC observations and predictions illustrates that ocean water influences groundwater SC in the southern part of the study area close to the Gulf of Mexico (pl. 1, callouts 3 and 6). Salt domes were not included as an explanatory variable and so were not tested as a source of salinity in the BRT models. The addition of geophysical explanatory variables added three-dimensional information about hydrogeology, including the location and thickness of confining units at the surface and with depth (James and Minsley, 2021; Minsley and others, 2021), that was unavailable to the baseline model. These additional variables allowed exploration of how surficial confinement or connection with deeper units may influence SC.

Effects of Deeper Saline Aquifers on Salinity in the Alluvial Aquifer

High SC may occur in the alluvial aquifer where there is a deeper source of salinity in connection with the alluvial aquifer (Kresse and Clark, 2008; Larsen and others, 2021). The bsmrva_25 geophysical explanatory variable is an estimation of the vertical hydraulic connectivity 25 m above and below the base of the alluvial aquifer; high values indicate poor hydraulic connection (James and Minsley, 2021) (fig. 12I). In contrast to conceptual models, the geophysical model predicted high SC where bsmrva_25 was high (fig. 12I), which corresponds to low mres and indicates that fine-grained sediments may be present (fig. 4). Broad areas of high predicted SC in southeastern Arkansas and northeastern Louisiana are coincident with the subcropping Vicksburg-Jackson confining unit (pl. 1). This area includes observations of high SC groundwater and good model performance in the Boeuf subregion (fig. 6). Specifically in Chicot County, Ark. (fig. 3), high SC occurs in an area where mres suggests that coarser grained sediments may be present and bsmrva_25 suggests that the alluvial aquifer may be connected with deeper, more saline aquifers; however, this pattern is not evident for the entire study area. The geophysical model may be predicting high SC where confining units are present because the resistivity signal captures both aquifer and porewater resistivity. For example,

in northeastern Arkansas, high SC predictions are coincident with an area where *bsmrva_25* is high (pl. 1, callouts 1 and 4); however, high SC observations are only localized around the Bald Knob, Ark., area (fig. 3). The geophysical model may be conflating the pattern of high SC and confining units in some areas of the model, thus causing overprediction of SC in other areas and increasing the difference in predictions between the two models (pl. 1).

Effects of Surficial Confining Characteristics on Salinity in the Alluvial Aquifer

The presence or absence of a surficial confining unit has been hypothesized to control salinity in some areas of the alluvial aquifer (Kresse and Clark, 2008). The absence, presence, and thickness of the surficial confining unit were represented by several geophysical explanatory variables, including *cnf_top*, *cnf_bot*, *cnf_class*, and *cnf_thk* (table 1, at back of report). As discussed in the “Discussion” section, *cnf_top* (fig. 12*X*) and *cnf_bot* (fig. 12*W*) are more representative of position in the groundwater system in the geophysical BRT model than is *cnf_class*. The geophysical model did not find *cnf_thk* to be an important predictor of SC (fig. 11). *Cnf_class*, a categorical variable, represents the degree of surficial confinement and was the 16th most influential explanatory variable. Low *cnf_class* categories correspond to fine-grained, low-permeability materials (James and Minsley, 2021). Low *cnf_class* categories were predicted to have higher SC than high *cnf_class* categories, though the variation was small (Killian and Knierim, 2023).

Relative Influences of Lithology and Salinity

Mres at the midpoint of the well screen was the seventh most influential explanatory variable in the geophysical model (fig. 11*B*). *Mres* captures both aquifer properties and porewater salinity, and previous analysis of the Mississippi Alluvial Plain region has shown that saline groundwater can influence resistivity measurements (Minsley and others, 2021). There is an inverse relation between SC and resistivity (fig. 14). Therefore, the geophysical model provides additional evidence that *mres* includes a resistivity signal driven, in part, by high SC (or saline) groundwater (fig. 14). *Mres* provides important three-dimensional information to the geophysical model and can be used to delineate contacts between the alluvial aquifer and deeper Tertiary confining units or aquifers (figs. 1, 4). However, because of the presence of saline groundwater in some areas of the alluvial aquifer (fig. 3), low *mres* (or high conductance) signals in the AEM survey data are not well constrained (Ward, 1990; Hilmi and Madun, 2022). Low resistivity can be the result of high salinity, the presence of fine-grained lithology, or a combination of the two (Zohdy and others, 1974). More work is needed to resolve the groundwater SC (salinity) signal within AEM-derived resistivity data across the Mississippi Alluvial Plain region. Removing the salinity

signal could allow for better estimation of subsurface materials, thus allowing for more accurate estimates of available water resources.

Lack of Temporal SC Variations at Wells

The year an SC measurement was collected was included as an explanatory variable in both models and had a similar influence on SC predictions (fig. 11). The SC dataset includes sampling networks from six States and includes long-term State cooperative programs to monitor salinity in Louisiana and Arkansas (Killian and Knierim, 2023), but network sampling priorities have changed over time, and there are gaps in the spatial and temporal coverage of data. For example, the Deltaic and Chenier Plains subregion shows elevated SC (fig. 3) but lacks data compared to other subregions (fig. 2*A*). Most SC measurements in this subregion were made prior to the year 1980 (fig. 2*A*). In contrast, there was an increase in SC observations after 1980 in the Delta subregion (fig. 2*A*). There may also be a bias introduced from networks that have the goal of monitoring localized high-salinity areas. The importance of this explanatory variable is likely driven by differences in sampling networks, as opposed to the BRT models capturing systematic changes in SC conditions throughout the aquifer.

In addition to the variation in the sampling network over time, groundwater quality may change at a location as water is withdrawn for irrigation or the aquifer is recharged (Kingsbury and others, 2014; Jurgens and others, 2020). Because only the most recent sample was used if a well had more than one observation, the model was not given information about how SC changed over time at a well. The SC maps were predicted using the year 2018 because 2018 was coincident with data collection for several explanatory variables. Although SC prediction maps could be made for any specified year, it would not be appropriate to use the maps to evaluate SC change through time across the aquifer. To understand how SC changes through time and which explanatory variables drive the change, explanatory variables that capture these changes through time are needed in addition to time-series SC observations.

Uncertainty in High SC in the Mississippi Alluvial Plain Region

The baseline and geophysical models showed overall similar trends in predicting broad areas of low SC in the northern part of the Mississippi Alluvial Plain region and high SC from southeastern Arkansas to the Gulf of Mexico (pl. 1). There are also several areas of localized known high SC in the Mississippi Alluvial Plain region (fig. 3), and both models were able to predict elevated SC in these areas (pl. 1). However, the spatial extent and magnitude of SC predictions were different between the baseline and geophysical models (pl. 1). Depending on which model is used, the same location

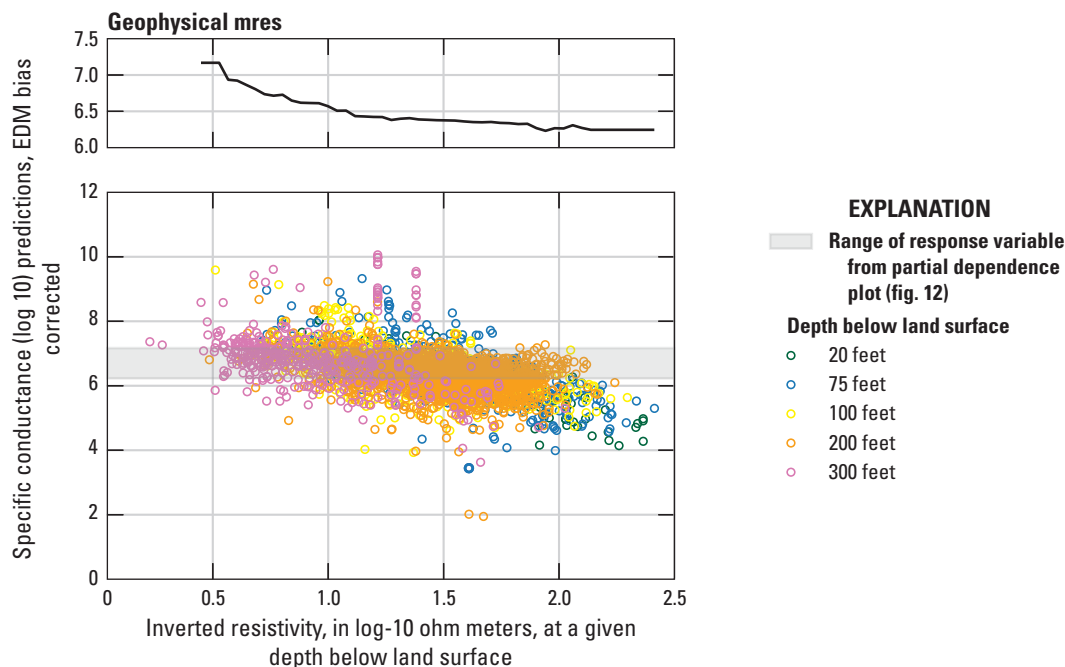


Figure 14. Comparison of inverted resistivity (mres) and specific conductance predictions for a given depth below land surface at groundwater wells in the Mississippi Alluvial Plain region by the geophysical model. See [table 1](#) (at back of report) for explanatory variable definitions and units; EDM, empirical distribution matching.

can be predicted with high confidence to have either fresh or brackish groundwater. For example, the Grand Prairie subregion ([fig. 1](#)) is predicted to have more brackish groundwater in the geophysical model than in the baseline model ([fig. 10](#)). Both models have similar overall accuracy ([fig. 5](#)) and similar accuracy among prediction depth ranges ([fig. 7](#)) and subregions ([fig. 6](#)). Therefore, both models have a similar magnitude of overpredicting or underpredicting SC across the Mississippi Alluvial Plain region ([fig. 5](#)).

The appropriate model and prediction maps to use depend on the goal and area for assessing SC conditions. Both models showed poor prediction accuracy for shallow SC conditions ([fig. 7](#)). SC predictions near the Gulf of Mexico from the geophysical model may not be realistic because of the structured nulls in training data ([fig. 6](#), pl. 1). Additionally, the high SC predictions in northeastern Arkansas from the geophysical model may be spurious because of the lack of high SC observations in that area. In contrast, the baseline model was trained with no three-dimensional data other than well depth, so the geophysical model may be more appropriate to use when making deeper SC predictions ([fig. 7](#)). Both the baseline and geophysical models can make accurate SC predictions for the Mississippi Alluvial Plain region with the given input data, but model outputs should be selected and interpreted on the basis of the scientific goal.

Summary

Groundwater is an important source of freshwater for agriculture and drinking water in the United States and around the world. Elevated specific conductance (SC) in aquifers can limit where freshwater is available for use. The Mississippi Alluvial Plain region is a major source of groundwater for irrigation in the south-central United States and includes areas of high SC that limit use. In 2021, the U.S. Geological Survey trained two boosted regression tree (BRT) machine-learning models on SC data available between 1942 and 2020 to predict groundwater SC across the Mississippi Alluvial Plain region at multiple depths below land surface for the Mississippi River Valley alluvial and deeper aquifers. Two BRT models were created, a baseline model and a geophysical model, to test and evaluate how the incorporation of aerial electromagnetic (AEM) data from an aquifer-wide geophysical survey affected model predictive performance and sensitivity. Geophysical data provided information about the three-dimensional hydro-geologic properties of aquifers and confining units, which affect groundwater flow paths and geochemistry.

The BRT models were developed using SC measurements collected at wells between 1942 and 2020, and predictions were made at a 1-kilometer resolution at multiple depths. Final models included explanatory variables that were more influential than random noise. The final baseline model was trained with 22 of the 29 available explanatory variables, which predominantly represented surficial information.

The final geophysical model incorporated geophysical data from a regional AEM survey and was trained on 33 of the 48 available explanatory variables. Twelve of the explanatory variables in the final geophysical model were from the AEM survey. Explanatory variables that describe position in the groundwater system were important drivers of SC for both models. Both models accurately predicted SC across the Mississippi Alluvial Plain region and corresponded to observed SC; the bias-corrected baseline model training coefficient of determination (R^2) value and holdout R^2 value were 0.896 and 0.617, and the bias-corrected geophysical model training R^2 value and holdout R^2 value were 0.906 and 0.626, respectively. Both models were able to predict elevated SC in known areas of high SC. The addition of AEM survey data did not improve the model globally, likely because of how effective BRT models are at identifying patterns in data. When the baseline and geophysical models are compared by region and depth, however, there is improvement in the geophysical model. AEM survey data provided additional three-dimensional information, whereas the baseline model relied on surficial variables and well depth to predict salinity at depth.

The models help to predict the extent of areas with elevated SC both vertically and horizontally for the Mississippi Alluvial Plain region. The baseline and geophysical models predicted similar areas of low SC in the northern part of the Mississippi Alluvial Plain region and high SC from southeastern Arkansas to the Gulf of Mexico. Geophysical data were available for the northern part of the study area and did not extend to the Gulf of Mexico. The missing geophysical explanatory variables for the southern part of the study area caused a linear artifact in SC predictions in the geophysical model. Therefore, BRT models can make predictions where explanatory variables are missing but cannot compensate for large areas of missing data (structured nulls). Because of the lack of geophysical explanatory variables for the southern part of the study area, SC predictions by the geophysical model are likely unreliable in that area. The geophysical data provided three-dimensional information that the baseline model did not have, which caused the geophysical model to predict greater SC with depth. Machine-learning models are useful for predicting groundwater quality in unmonitored locations at the regional scale, but model limitations and prediction uncertainty should be considered within the context appropriate for the system and scientific goals.

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Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.

[NA, not applicable; cm, centimeter; mm, millimeter; m, meter; +/-, above and below; AEM, aerial electromagnetic; NAD 83, North American Datum of 1983; mm/dd/yyyy hh:mm, month/day/year hour:minute; NAVD 88, North American Vertical Datum of 1988; xml, Extensible Markup Language; km, kilometer; yyyy, year; EDM, empirical distribution matching]

Name	Units or format	Description	Final base-line model variable	Final geophysical model variable	Dataset group	Citation
SITEAG	NA	Well identifier	NA	NA	Supplementary information	U.S. Geological Survey (2021)
AQFR_CD	eight-character string	Aquifer code (from https://help.waterdata.usgs.gov/aqfr_cd#:~:text=Aquifer%20Codes%20(aqfr_cd),by%20the%20Water%20Resources%20Division.%22)	NA	NA	Supplementary information	U.S. Geological Survey (2021)
AWC_0200	inches per foot	Available water capacity (AWC), quantity of water soil can store for use by plants, 0 to 200 cm depth	No	No	Input, explanatory variable, soils	Soil Survey Staff (2020)
BD_0200	grams per cubic centimeter	Bulk density (BD), oven-dried weight of the soil material less than 2 mm in size per unit volume of soil at water tension of 1/3 bar, 0 to 200 cm depth	Yes	Yes	Input, explanatory variable, soils	Soil Survey Staff (2020)
bsmrva_5	5 m +/- the Mississippi River Valley alluvial aquifer base	Vertically integrated electrical conductance values within the vertical interval spanning +/- 5 m across the base of the Mississippi River Valley alluvial aquifer	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
bsmrva_10	10 m +/- the Mississippi River Valley alluvial aquifer base	Vertically integrated electrical conductance values within the vertical interval spanning +/- 10 m across the base of the Mississippi River Valley alluvial aquifer	NA	No	Input, explanatory variable, AEM	James and Minsley (2021)
bsmrva_25	25 m +/- the Mississippi River Valley alluvial aquifer base	Vertically integrated electrical conductance values within the vertical interval spanning +/- 25 m across the base of the Mississippi River Valley alluvial aquifer	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
BOT	feet	Bottom of well screen	Yes	Yes	Input, explanatory variable, well construction	U.S. Geological Survey (2021)
clay_0200	percentage	Clay, soil separate consists of mineral soil particles that are less than 0.002 mm in diameter, 0 to 200 cm depth	No	No	Input, explanatory variable, soils	Soil Survey Staff (2020)

Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.—Continued

[NA, not applicable; cm, centimeter; mm, millimeter; m, meter; +/-, above and below; AEM, aerial electromagnetic; NAD 83, North American Datum of 1983; mm/dd/yyyy hh:mm, month/day/year hour:minute; NAVD 88, North American Vertical Datum of 1988; xml, Extensible Markup Language; km, kilometer; yyyy, year; EDM, empirical distribution matching]

Name	Units or format	Description	Final base- line model variable	Final geophys- ical model variable	Dataset group	Citation
cnf_bot	meters	Bottom altitude corresponding to the surficial confining material of the Mississippi River Valley alluvial aquifer	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
cnf_class	category -3 to 3	Metric representing the degree of surface confinement or connectivity that ranges from fully confining conditions to high potential hydrologic connectivity of the Mississippi River Valley alluvial aquifer	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
cnf_thk	meters	Occurrence and thickness of surficial (less than 15 m depth) confining material of the Mississippi River Valley alluvial aquifer	NA	No	Input, explanatory variable, AEM	James and Minsley (2021)
cnf_top	meters	Top altitude corresponding to the surficial confining material of the Mississippi River Valley alluvial aquifer	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
d2w_ft_bls	feet below land surface	Depth to water, spring 2018, Mississippi River Valley alluvial aquifer	Yes	Yes	Input, explanatory variable, hydrology	McGuire and others (2023)
DEC_LAT_VA	decimal degrees, NAD 83	Latitude	Yes	Yes	Input, explanatory variable, location information	U.S. Geological Survey (2021)
DEC_LONG_V	decimal degrees, NAD 83	Longitude	Yes	Yes	Input, explanatory variable, location information	U.S. Geological Survey (2021)
dtm	mm/dd/yyyy hh:mm	Date and time stamp of water-quality sample collection	NA	NA	Supplementary information	U.S. Geological Survey (2021)
ER_0013	inches per year	Empirical water balance (EWB) model, estimates of effective recharge, years 2000–13	Yes	Yes	Input, explanatory variable, climate	Reitz and others (2017a, b)
ET_0013	inches per year	Empirical water balance (EWB) model, estimates of evapotranspiration, years 2000–13	Yes	Yes	Input, explanatory variable, climate	Reitz and others (2017a, b)

Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.—Continued

[NA, not applicable; cm, centimeter; mm, millimeter; m, meter; +/–, above and below; AEM, aerial electromagnetic; NAD 83, North American Datum of 1983; mm/dd/yyyy hh:mm, month/day/year hour:minute; NAVD 88, North American Vertical Datum of 1988; xml, Extensible Markup Language; km, kilometer; yyyy, year; EDM, empirical distribution matching]

Name	Units or format	Description	Final base-line model variable	Final geophysical model variable	Dataset group	Citation
facies_ <i>[d]</i>	class 0 to 20	Facies class for materials expected to have similar hydrologic and geologic properties based on electrical resistivity for the depth (d) of interest. For well points, facies class is calculated for the screened interval	NA	No	Input, explanatory variable, AEM	James and Minsley (2021)
Formation	formation symbol; see Wacaster and others (2018) for explanation	Saucier geomorphology by formation	Yes	Yes	Input, explanatory variable, geology	Wacaster and others (2018)
GeoAge	category	Saucier geomorphology formation grouped by geologic age (Holocene or Pleistocene)	No	No	Input, explanatory variable, geology	Wacaster and others (2018)
GW2018_ft	feet, NAVD 88	Potentiometric surface of the Mississippi River Valley alluvial aquifer for the year 2018	Yes	Yes	Input, explanatory variable, hydrology	McGuire and others (2020a, b)
HSG_100m	character string	Hydrologic soil groups (HSG), classes that indicate runoff potential for soils. See Killian and Knierim (2023) for category description.	No	No	Input, explanatory variable, soils	Soil Survey Staff (2020)
Ksat_0200	inches per day	Saturated hydraulic conductivity (Ksat), ability of a saturated soil to transmit water, 0 to 200 cm depth	Yes	Yes	Input, explanatory variable, soils	Soil Survey Staff (2020)
Layer	character string	Flag for group of aquifer codes. See Killian and Knierim (2023) for category description.	NA	NA	Supplementary information	NA
ls_alt	feet	Land surface altitude, used to calculate depth of prediction for rasters	No	No	Supplementary information	U.S. Geological Survey (2018)
magint	nano Teslas	Residual magnetic intensity	No	Yes	Input, explanatory variable, AEM	Burton and others (2021)
MOHP_ <i>[i]</i> DSD	meters	Multi-order hydrologic position (MOHP) of ith-order streams, distance to stream divide (DSD)	9, 5, 4	9, 5, 4	Input, explanatory variable, hydrology	Moore and others (2019), Belitz and others (2019)
MOHP_ <i>[i]</i> LP	meters	Multi-order hydrologic position (MOHP) of ith-order streams, lateral position (LP)	9, 5, 4	9, 5, 4	Input, explanatory variable, hydrology	Moore and others (2019), Belitz and others (2019)

Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.—Continued

[NA, not applicable; cm, centimeter; mm, millimeter; m, meter; +/–, above and below; AEM, aerial electromagnetic; NAD 83, North American Datum of 1983; mm/dd/yyyy hh:mm, month/day/year hour:minute; NAVD 88, North American Vertical Datum of 1988; xml, Extensible Markup Language; km, kilometer; yyyy, year; EDM, empirical distribution matching]

Name	Units or format	Description	Final base-line model variable	Final geophysical model variable	Dataset group	Citation
mres_ <i>[d]</i>	log 10	Inverted resistivity (ohm meters) at a given depth (<i>d</i>) below land surface	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021), Minsley and others (2021)
mrvaBasElv	meters	Mississippi River Valley alluvial aquifer base altitude	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
MRVAbtmalt	feet	Bottom of the Mississippi River Valley alluvial aquifer, altitude	Yes	Yes	Input, explanatory variable, hydrogeologic framework	Torak and Painter (2019a, b)
mrvaSatThk	meters	Saturated Mississippi River Valley alluvial aquifer thickness	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
mrvaTopElv	meters	Top altitude of the saturated Mississippi River Valley alluvial aquifer based on the 2018 potentiometric surface and digital elevation model	NA	No	Input, explanatory variable, AEM	James and Minsley (2021)
NADP_cl	milligrams per liter	National Atmospheric Deposition Program (NADP) 26-year average chloride (Cl) concentration	No	No	Input, explanatory variable, climate	National Atmospheric Deposition Program (2021)
NADP_na	kilograms per hectare	National Atmospheric Deposition Program (NADP) 26-year average sodium (Na) concentration	Yes	No	Input, explanatory variable, climate	National Atmospheric Deposition Program (2021)
NAT_AQFR_C	character string	National aquifer code, reference list available at https://www.usgs.gov/mission-areas/water-resources/science/national-aquifer-code-reference-list	NA	NA	Supplementary information	U.S. Geological Survey (2021)
noise	NA	Random noise	No	No	Input, explanatory variable	NA
QF_0013	inches per year	Empirical water balance (EWB) model, estimates of quickflow (runoff), years 2000–13	Yes	Yes	Input, explanatory variable, climate	Reitz and others (2017a, b)
r00095	less than and greater than the specific conductance value	Flag for censored values	NA	NA	Supplementary information	U.S. Geological Survey (2021)

Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.—Continued

[NA, not applicable; cm, centimeter; mm, millimeter; m, meter; +/-, above and below; AEM, aerial electromagnetic; NAD 83, North American Datum of 1983; mm/dd/yyyy hh:mm, month/day/year hour:minute; NAVD 88, North American Vertical Datum of 1988; xml, Extensible Markup Language; km, kilometer; yyyy, year; EDM, empirical distribution matching]

Name	Units or format	Description	Final base-line model variable	Final geophysical model variable	Dataset group	Citation
region	see Mississippi Alluvial Plain generalied_regions_FGDC_metadata.xml available at the citation for category description	Seven generalized subregions of the Painter and Westerman (2018) extent of the Mississippi Alluvial Plain to provide a framework for analysis, visualization, and regional comparisons of local data within the region.	Yes	Yes	Input, explanatory variable, location information	Ladd and Travers (2019)
sand_0200	percentage	Sand, soil separate consists of mineral soil particles that are 0.05 to 2 mm in diameter, 0 to 200 cm depth	Yes	Yes	Input, explanatory variable, soils	Soil Survey Staff (2020)
SC	microsiemens per centimeter at 25 degrees Celsius	Specific conductance value	NA	NA	Input, response variable	U.S. Geological Survey (2021)
silt_0200	percentage	Silt, soil separate consists of mineral soil particles that are 0.002 to 0.05 mm in diameter, 0 to 200 cm depth	No	No	Input, explanatory variable, soils	Soil Survey Staff (2020)
soilrd_K	percentage	Potassium soil radiometrics	NA	No	Input, explanatory variable, AEM	Burton and others (2021)
soilrd_Th	parts per million	Thorium (eTh) soil radiometrics	NA	Yes	Input, explanatory variable, AEM	Burton and others (2021)
soilrd_ThK	ratio	Soil radiometric ratio of thorium and potassium (rThK)	NA	Yes	Input, explanatory variable, AEM	Burton and others (2021)
soilrd_U	parts per million	Uranium (eU) soil radiometrics	NA	No	Input, explanatory variable, AEM	Burton and others (2021)
soilrd_UK	ratio	Soil radiometric ratio of uranium and potassium (rUK)	NA	No	Input, explanatory variable, AEM	Burton and others (2021)
SOURCE	category	Source of well construction information and water-quality data	NA	NA	Supplementary information	NA

Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.—Continued

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Name	Units or format	Description	Final base- line model variable	Final geophys- ical model variable	Dataset group	Citation
StrmCon	siemens	Streambed connectivity, vertically integrated electrical conductance values from 0 to 10 m depth for all locations within 1 km radius of streams within the Mississippi Alluvial Plain region	NA	Yes	Input, explanatory variable, AEM	James and Minsley (2021)
subUnit	category	Subcropping hydrogeologic unit	NA	NA	Supplementary information	Hart and others (2008)
TOP	feet	Top of well screen	NA	NA	Supplementary information	U.S. Geological Survey (2021)
WELL_DEPTH	feet	Depth of well below land surface	NA	NA	Supplementary information	U.S. Geological Survey (2021)
year	yyyy	Year sample was collected	Yes	Yes	Input, explanatory variable	U.S. Geological Survey (2021)
m1Pred_In	natural log	Baseline model predicted specific conductance	NA	NA	Output, prediction	Killian and Knierim (2023)
m1P_uScm	microsiemens per centimeter	Baseline model predicted specific conductance	NA	NA	Output, prediction	Killian and Knierim (2023)
m1edmP_In	natural log	Predicted specific conductance with EDM bias correction for the baseline model	NA	NA	Output, prediction	Killian and Knierim (2023)
m1edmP_uScm	microsiemens per centimeter	Baseline model EDM bias correction for predicted specific conductance, backtransformed to original units	NA	NA	Output, prediction	Killian and Knierim (2023)
m1res_uScm	microsiemens per centimeter	Baseline model residuals (observed – predicted specific conductance)	NA	NA	Output, prediction	Killian and Knierim (2023)
m1rEDM_uScm	microsiemens per centimeter	Baseline model EDM bias correction of residuals (observed – predicted specific conductance)	NA	NA	Output, prediction	Killian and Knierim (2023)
m2Pred_In	natural log	Geophysical model predicted specific conductance	NA	NA	Output, prediction	Killian and Knierim (2023)
m2P_uScm	microsiemens per centimeter	Geophysical model predicted specific conductance	NA	NA	Output, prediction	Killian and Knierim (2023)
m2edmP_In	natural log	Predicted specific conductance with EDM bias correction for the geophysical model	NA	NA	Output, prediction	Killian and Knierim (2023)

Table 1. Name, unit, description, dataset group, and source information for each variable in the baseline and geophysical models to predict groundwater specific conductance across the Mississippi Alluvial Plain region.—Continued

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Name	Units or format	Description	Final base-line model variable	Final geophysi-cal model variable	Dataset group	Citation
m2edmP_uScm	microsiemens per centimeter	Geophysical model EDM bias correction of pre-dicted specific conductance, backtransformed to original units	NA	NA	Output, predic-tion	Killian and Knierim (2023)
m2res_uScm	microsiemens per centimeter	Geophysical model residuals (observed – predicted specific conductance)	NA	NA	Output, predic-tion	Killian and Knierim (2023)
m2rEDM_uScm	microsiemens per centimeter	Geophysical model EDM bias correction of residu-als (observed – predicted specific conductance)	NA	NA	Output, predic-tion	Killian and Knierim (2023)

For more information about this publication, contact
Director, [Lower Mississippi-Gulf Water Science Center](#)
U.S. Geological Survey
640 Grassmere Park, suite 100
Nashville, TN 37211

For additional information, visit
<https://www.usgs.gov/centers/lmg-water/>

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