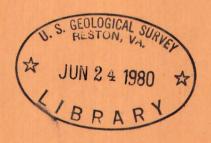
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Geology of the Quartz Creek Pegmatite District, Gunnison County, Colorado



Trace Elements Investigations Report 138

UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

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Geology - Mineralogy

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Series A

UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

GEOLOGY OF THE QUARTZ CREEK PEGMATITE DISTRICT,
GUNNISON COUNTY, COLORADO*

By

Mortimer H. Staatz and Albert F. Trites

April 1952

Trace Elements Investigations Report 138

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*This report concerns work done on behalf of the Division of Raw Materials of the U. S. Atomic Energy Commission

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GEOLOGY OF THE QUARTZ CREEK PEGMATITE DISTRICT, GUNNISON COUNTY, COLORADO

by

Mortimer H. Staatz and Albert F. Trites

ABSTRACT

The Quartz Creek pegmatite district includes an area about 29 square miles in the vicinity of Quartz Creek in Gunnison County, Colo. This area contains, 1,803 pegmatites that are intruded into pre-Cambrian rocks.

The rocks exposed in the district range in age from pre-Cambrian to Recent. The oldest pre-Cambrian rocks are chiefly quartzites interbedded with a few arkoses and conglomerates. These rocks are surrounded by more abundant hornblende gneiss and tonalite. A small body of biotite tonalite was intruded and two thin layers of dacitic pillow lava were extruded into this series. The hornblende gneiss and tonalite have the same composition and differ only in texture. The older material (hornblende gneiss) has a well-marked lineation, whereas the younger (tonalite) is equigranular. Subsequently, a large body of quartz monzonite was intruded along the northern boundary of the mapped area. Later, coarse-grained granite was intruded into the southern part of the area. Dikes of fine-grained granite cut the coarse-grained variety. The last period of intrusive activity in pre-Cambrian time is marked by a large number of pegmatites.

The pre-Cambrian rocks were tilted and eroded, and the flatlying

Jurassic Morrison formation was deposited on the irregular surface. This

formation is conformably overlain by the Cretaceous Dakota sandstone.

Faulting produced a vertical offset of 410 feet in the Mesozoic sediments

along the only large fault in the area. At the end of Mesozoic time there

was another period of erosion. Tertiary (?) tuff is exposed in small, scattered areas in the southern part of the district. It overlies both the Dakota sandstone and pre-Cambrian formations. Glacial till occurs along the edges of Quartz Creek and Wood Gulch. Quaternary alluvium fills the valley bottoms.

Although the composition of the country rock has little effect on the shape of the pegmatites, the foliation imposed on this rock has a localizing effect and in part controls the ultimate shape of pegmatites. Zoned and related internal structures are not well developed in the pegmatites of this region. Many of the pegmatites are homogeneous and those that are zoned usually contain a large wall zone and small discontinuous cores. In addition to the more common homogeneous and zoned pegmatites, 7 percent of the pegmatites show a layered structure of textural and mineralogical units not repeated on the opposite side of the pegmatite. Other internal structural units include pegmatites which vary in composition along strike, multiple or "line-rock pegmatites", and fracture fillings.

The mineralogy of the pegmatites is described in detail. Specific attention was given to most of the 27 observed minerals. A study of the index of refraction of 439 specimens of plagioclase showed that the variation from zone to zone and layer to layer is minor and that there is no systematic variation in respect to the entire district. No correlation could be found between the refractive index of plagioclase in the pegmatites and the type of country rock, or the presence of various accessory minerals.

Index of refraction determinations on 95 specimens of muscovite showed no constant variation from wall zone to core or from layer to layer. Curved muscovite has identical optical properties with the flat variety.

The index of refraction was determined for 183 beryl specimens. The beryl in pegmatites containing only a wall zone and a core showed no difference between zones, but in pegmatites that have intermediate zones, the indices of refraction of the beryl showed an inward increase in the alkali content from the contact. Beryl occurs with almost all of the pegmatite minerals and is not restricted in its mineral associations.

Tourmaline, except the black variety, is associated with lepidolite.

Dark green and blue tourmaline is found in the outer zones of pegmatites containing lepidolite, and the pink and light green varieties are found in direct contact with lepidolite.

Lepidolite occurs in aggregates of fine grains, in flat plates, and in curved plates; the three varieties are optically identical. The lighter-colored varieties have higher indices of refraction and contain less lithia than the darker varieties.

In addition, the occurrence of the following minerals is described in detail: perthite, quartz, martite, biotite, garnet, columbite-tantalite, monazite, microlite, topaz, gahnite, allanite, and an unidentified mineral.

The lack of alteration in the wall rocks adjacent to the pegmatites is interpreted as indicating that the original pegmatite magma did not have an excess of materials such as B, OH-, and P that are needed to form alteration minerals. Because of their low concentration, the above materials were available only in the pegmatitic magma during its crystallization. Pegmatites that contain the rare minerals such as beryl, tourmaline, curved muscovite, biotite, magnetite, monazite, columbite-tantalite, cleavelandite, topaz, lepidolite, and microlite show a grouping in clusters within the district.

Beryl-bearing pegmatites occur most abundantly in hornblende gneiss and are only rarely found in either granite or quartz monzonite.

The types of minerals that form in a pegmatite appear to be determined by the character of the material segregated from the original magma and the period in which it segregated. The elements escape at one period and may be from only a specific pocket in the magma. These liquids tend to form groups of pegmatites in which the later bodies contain a high proportion of volatiles.

Inferred reserves of the district are estimated for beryl, scrap mica, both hand-cobbing and milling feldspar, lepidolite, columbite-tantalite, topaz, monazite, and microlite. No sheet mica was found. Reserves are small and transportation costs are high so substantial production of low-priced feldspar and scrap mica will depend on the adoption of economical milling techniques for recovering the large quantities of feldspar available. Beryl is irregularly distributed and its recovery as a byproduct will depend on the establishment of a stable market for feldspar and scrap mica. Lepidolite reserves are small and low grade.

INTRODUCTION

Prior to World War II the Quartz Creek pegmatite district was well known for its fine specimens of colored tourmaline, books of lepidolite, topaz, and microlite. During and after World War II small quantities of beryl, feldspar, lepidolite, and tantalum minerals were produced from the district.

Location and surface features

The Quartz Creek pegmatite district is on the western slope of the Sawatch Range in Tps. 49 and 50 N., R. 3 E., New Mexico principal meridian, Gunnison County, Colo. (fig. 1). It covers about 29 square miles in the

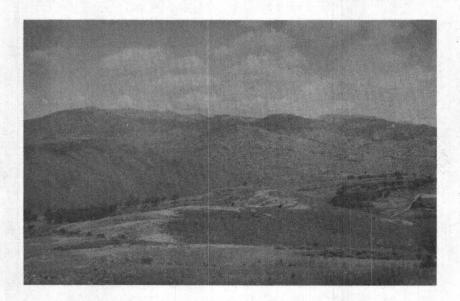
vicinity of Quartz Creek. State Highway 162 follows Quartz Creek through the district, and joins U. S. Highway 50, 2 miles south of the southern boundary of the mapped area. A road branching from State Highway 162 crosses the southeastern corner of the area and connects with U. S. Highway 50 mile west of Doyleville. Unimproved roads follow several of the valleys, and a mine-access road, made during World War II, connects State Highway 162 with the Brown Derby mine. The nearest railroad shipping point, Parlin, is on a narrow-gauge line of the Denver and Rio Grande Western Railroad, which connects with standard-gauge lines at Salida (east-bound) or Montrose (west-bound).

The topography is moderately steep and has a maximum relief of 2,200 feet (pl. I). The highest peak in the district is at the northern boundary and has an elevation of 10,238 feet. Quartz Creek, the main tributary draining the area, is in a flat cultivated valley, a quarter to half a mile wide. The hills rising from Quartz Creek are usually sage-covered, and the north-facing slopes of the higher hills are covered with aspen, or spruce and pine. Quartz and Alder Creeks are the only permanent streams in the district, but water flows in Willow Creek and Wood Gulch during the spring and early summer.

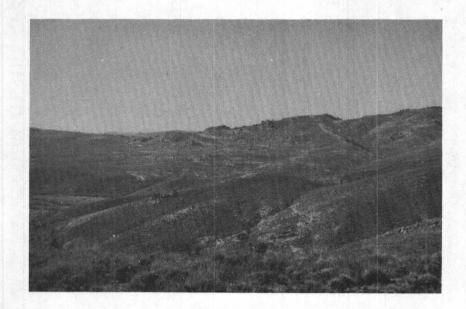
Production

The recorded production of pegmatite minerals from the Quartz Creek district is about 51 tons of beryl, 283 tons of lepidolite, 140 tons of scrap mica, 5,000 pounds of tantalum and columbium minerals, and 15 pounds of monazite.

From September 1943 to the spring of 1945 the Brown Derby property was leased by the Hayden Mining Company. Prior to February 1945, 3,155.67



A. Panorama southern end of Quartz Creek pegmatite district looking north. White areas in foreground are pegmatites.



B. Pancrama of the northern end of the Quartz Creek pegmatite district looking west across Quartz Creek. Outcrops in the background are the Black Wonder (No. 847) pegmatite.

pounds _/ of beryl were sold to Metals Reserve Company and 283 tons of

✓ Hanley, J. B., unpublished notes.

lepidolite to Corning Glass Company. In addition, 4,000 pounds of microlite concentrate containing 52 percent microlite was stockpiled at the mill and later sold. Though the final production figures are not available, they do not greatly exceed these figures as mining stopped in the spring of 1945.

The White Spar No. 1 and No. 2 pegmatites, which are 0.8 mile north of the Brown Derby mine, were mined for a short time during World War II by the Colorado Feldspar Company. The production of lepidolite and feldspar is not known, but was small.

There was no mining in the district from 1945 to 1947. Mr. Rod

Fields located the Bucky claim on the east side of Willow Creek and started

to mine beryl in the spring of 1948. Mr. Fields produced 17 tons of beryl,

and in November 1948 sold the property to Beryllium Mining Company, Inc.,

which has produced 32.0 tons of beryl, 139.6 tons of scrap mica, 1,020

pounds of columbite-tantalite, 15 pounds of monazite, and 13 pounds of a

samarskite-like mineral to May 1950. The last two minerals were sold to

Ward's Natural Science Establishment, Inc.

Previous work

Early papers on the pegmatites of the Quartz Creek district have been concerned chiefly with the mineralogy of the Brown Derby pegmatites \mathcal{L} . A

____ Eckel, E. B., A new lepidolite deposit in Colorado: Am. Ceramic Soc. Jour., vol. 16, pp. 239-245, 1933.

Landes, K. K., Colorado pegmatites: Am. Mineralogist, vol. 20, p. 333, 1935.

Eckel, E. B., and Lovering, T. S., Work of Eckel, Lovering, Fair-child-Microlite from Ohio City, Colorado: Report of the Committee on the Measurement of Geologic Time, pp. 77-79, Apr. 1, 1935.

Seaman, D. M., New pegmatite locality near Ohio City, Colorado: Oregon Mineralogist, vol. 2, p. 23, 1934.

map showing the regional geology of the Gold Brick district, on a scale of 1.5 inches equal one mile, was published by Crawford and Worcester _____. The

_/ Crawford, R. D., and Worcester, P. G., Geology and ore deposits of the Gold Brick district, Colorado: Colorado Geol. Survey Bull. 10, 1916.

southwestern corner of their map, an area roughly 3.3 miles by 2 miles, overlaps the northern part of the present area mapped. No pegmatites are shown on their map and the area containing them is designated as "granite". The area around Tomichi Dome, several miles to the east of the Quartz Creek district, has been described by Stark and Behre _/.

_/ Stark, J. T., and Behre, C. H., Jr., Tomichi dome flow: Geol. Soc. America Bull., vol. 47, pp. 101-110, 1936.

Between September 1942 and December 1944 the Geological Survey had several field parties mapping in Colorado under E. W. Heinrich in 1942 and John B. Hanley in 1942-1944 J. In the Quartz Creek district these parties

mapped in detail—on scales of 1 inch equals 20 feet to 1 inch equals 50 feet—the Opportunity No. 1 claim, the Brown Derby No. 1 claim, the Brown Derby Ridge pegmatites, Brown Derby No. 5, the White Spar No. 1, the White Spar No. 2, and the Bazooka pegmatites. A total of 25 pegmatites was mapped with plane table and telescopic alidade. Several other pegmatites were visited and described.

Field work and acknowledgments

The investigations carried out by the Geological Survey in the Quartz Creek pegmatite district during World War II were concerned primarily with pegmatites from which feldspar, muscovite, and minerals containing beryllium, tantalum, lithium, and the rare earths were produced. Such pegmatites are in the minority, and time did not permit detailed study of the more numerous associated, but unproductive pegmatites, or of the broader relations of pegmatites to the regional geology. This study, started in 1948, was made not only to provide an economic appraisal of individual deposits, but also to determine the regional relationships of pegmatites and country rock.

The field work for this report was started on July 10, 1948. M. H. Staatz and P. T. Flawn began mapping on the east side of Quartz Creek and A. F. Trites and F. L. Klinger on the west side. Field work was recessed September 7, 1948. It was resumed on June 12, 1949 and completed December 10, 1949. The writers were assisted during 1949 by F. L. Klinger for three months, and J. D. Vogel for two months. Mapping was done by pace and Brunton compass, using the Pitkin quadrangle topographic map enlarged to a scale of one inch equals 1,000 feet as a base (pl. 11). Individual pace and compass maps also were made of each pegmatite on the scales of from one inch equals 40 feet to one inch equals 200 feet, depending upon the size of the pegmatite. Petrographic work was done during the spring of 1950.

This investigation was made in part on behalf of the Division of Raw Materials of the U. S. Atomic Energy Commission.

GENERAL GEOLOGY

The rock units mapped in the Quartz Creek pegmatite district range in

age from pre-Cambrian to Recent (pl. LT). The age of the Brown Derby No. 1 pegmatite, as determined from uranium-bearing microlite by Eckel and Lovering _/, is 760 million years. The oldest pre-Cambrian rocks consist of

∠ Eckel, E. B., and Lovering, T. S., op. cit., p. 79, 1935.

metasedimentary rocks, predominantly quartzites, that are surrounded by younger and more abundant tonalite and hornblende gneiss (a meta-tonalite). Included in this series are two small bands of dacitic pillow lava and one of biotite tonalite. A coarse-grained porphyritic granite, similar in appearance _/ to the Pikes Peak granite, intrudes the earlier pre-Cambrian

✓ Eckel, E. B., op. cit., p. 240, 1933.

rocks in the south-central part of the district and a large quartz monzonite intrusive body occurs in the extreme northern part. These rocks are intruded by numerous fine-grained pink granite dikes and by a large number of pegmatites.

An angular unconformity separates the pre-Cambrian rocks from the flatlying Jurassic Morrison formation and Cretaceous Dakota sandstone which crop out along the east and west sides of the area.

Flatlying Tertiary (?) tuff is exposed in three scattered patches overlying unconformably both the Dakota sandstone and pre-Cambrian formations. Small areas of glacial till border Quartz Creek and Wood Gulch, and Quaternary alluvium fills many of the valley bottoms.

In general, the pre-Cambrian formations dip steeply and have a north-westerly trend, which is brought out by the structure of both the metasedimentary rocks, the dacitic pillow lava, and the biotite tonalite. The pegmatites have a general northeast trend across all of the earlier structures.

Only three faults with displacements of over 20 feet were found in the area. The largest of these trends northwest and separates the Dakota and Morrison formations from the pre-Cambrian in the southwestern corner of the district. Two other faults, which are terminated by this large fault, separate a block of Dakota sandstone from the pre-Cambrian and Morrison formations.

Pre-Cambrian rocks

Quartzite

Pre-Cambrian quartzite, interbedded partly with arkosic and conglomerate quartzite (pl. II), is best exposed on the northwestern slope of Wood Gulch where there are four mappable bodies. Two other bodies of arkosic quartzite, about half a mile long, crop out along the headwaters of Tollgate Gulch, a tributary of Quartz Creek, which is northwest of Wood Gulch. Narrow outcrops, a few tens of hundreds of feet long, are found at widely scattered localities on the northern side of Quartz Creek. These rocks have been highly metamorphosed and are part of a much larger area of sedimentary rocks that are separated by intrusive tonalite and hornblende gneiss.

The pre-Cambrian quartzites are generally dark gray but in places are white and brown. The original sediments ranged from siltstone to conglomerate but most were fine grained. Some of the quartzites are now slightly schistose. The northernmost band of metasedimentary rocks in Wood Gulch is a conglomerate containing pebbles from 0.1 to 2 inches long. Some of the pebbles are elongated; the ratio of width to length is from 1:4 to 1:5.

Feldspar (orthoclase, microcline, and plagioclase) is present throughout the unit but the proportion varies widely. The rocks along Wood Gulch are

Ash and the Wall of the

commonly quartzites with only a few percent of feldspar, but those on the north side of Quartz Creek contained 20 percent or more of feldspar. The proportion of the dominant dark mineral, biotite, ranges from a trace to about 15 percent. Muscovite is common in amounts of one percent or less. In rocks rich in feldspar, epidote is prominent and may make up more than 50 percent of the rock. One specimen contains hornblende and clinozoisite as well as biotite. Apatite, zircon, and magnetite are common accessories.

Quartz-mica schists, composed chiefly of quartz, biotite, feldspar, and muscovite, are found in a few scattered outcrops in the northern part of the area. Locally, these schists contain well-developed porphyroblasts of quartz and magnetite.

The thicknesses of the different exposures range from a few feet to a maximum of about 600 feet,

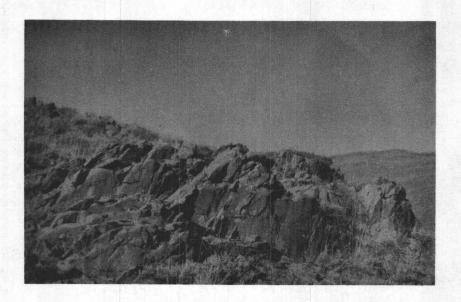
The quartzites are the oldest rocks in the district and are surrounded by younger hornblende gneiss, tonalite, and granite. One xenolith of conglomerate was found in the granite.

Dacite

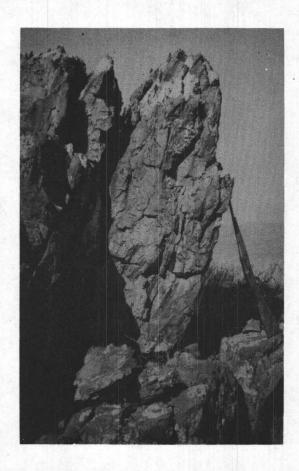
Dacitic pillow lava (pl. II) occurs in two northwesterly-trending bands south of Quartz Creek. One of these is in sec. 11, T. 50 N., R. 3 E. on the northwestern slope of Wood Gulch, and the other is in secs. 3 and 34, about 900 feet northwest of the Brown Derby mine.

The pillow lava is yellow green to dark green depending on the proportion of epidote. It is a fine-grained dense vesicular rock, with prominent flow lines. Some vesicles contain well-developed crystals of epidote and quartz; a few are completely filled with fine-grained quartz.

Large ellipsoids or pillows (pl. III, B), several feet long and about one



A. Tonalite outcrop on the northwest side of Wood Gulch.



B. Pillow lava showing ellipsoidal pillows on the northwest side of Wood Gulch.

foot wide, are common. Microscopically the unaltered rock consists of green prismatic hornblende (50 percent), quartz (30 percent), and andesine (20 percent). Epidote may be present almost to the exclusion of other minerals and is veined by calcite.

The band of pillow lava on the northwestern slope of Wood Gulch is 140 feet thick and is parallel in strike and dip to the enclosing pre-Cambrian quartzite and conglomerate. This pillow lava was extruded under water onto a sand and was buried by later sediments. A second band of pillow lava, 6,450 feet to the northwest, is enclosed in hornblende gneiss. The two areas of pillow lava (pl. II) are almost alined on strike and are probably remnants of the same band. Both bands of pillow lava are younger than the pre-Cambrian quartzites on the northeast and older than the quartzites to the southwest.

Biotite tonalite

One small body of biotite tonalite, approximately 900 feet long and 110 feet wide, is exposed in sec. 11 on the northwestern slope of Wood Gulch. It is bounded on the north by pre-Cambrian quartzite and on the south by hornblende gneiss, and is approximately 80 feet southwest of, and parallel to, a band of pillow lava.

The biotite tonalite is dark gray with prominent black hornblende crystals, 0.05 inch in diameter, in a black, speckled, fine-grained matrix. Plagioclase phenocrysts, the same size as the hornblende crystals, contain small included grains of biotite, epidote, and hornblende. Dark minerals make up about 50 percent of the rock. Hornblende, the chief dark mineral, constitutes 25 percent of the rock, and commonly forms ragged prismatic grains, but locally it occurs as small included grains in the plagioclase

and quartz. Biotite constitutes 22 percent of the rock and occurs with hornblende as aggregates and in the plagioclase as a myriad of fine grains. Epidote (3 percent) and magnetite (<1 percent) are the other dark minerals. Andesine (40 percent) forms large crystals clouded with many fine crystals of biotite, hornblende, and epidote. Quartz (10 percent) is interstitial to the andesine.

The biotite tonalite occurs near the southwest edge of the pre-Cambrian quartzite and has a trend parallel to the strike of the bedding. It is similar in composition to the pillow lava, except that it is not vesicular and contains much less epidote and no calcite. It is also much coarser grained. The similarity in trend and composition suggests that the biotite tonalite and the pillow lava derived from the same magma at about the same time, but one was intrusive and the other extrusive. The biotite tonalite is younger than the quartzite to the north and probably older than the hornblende gneiss to the south.

Hornblende gneiss and tonalite

The hormblende gneiss and tonalite are gradational—sometimes in the same outcrop. The difference between the two rocks is one of texture, with the gneiss exhibiting a planar alimement of minerals. Though most of the hormblende gneiss and tonalite is clearly of intrusive origin, some units are lava flows. The hormblende gneiss and tonalite were mapped separately only along the northwestern slope of Wood Gulch, where fine-grained hormblende gneiss is very schistose and is cut by the coarser, equigranular tonalite (pl. III, A). These two rocks are evidently of different ages, the tonalite having been intruded after the older rock had been metamorphosed considerably. The intrusive tonalite bodies in Wood Gulch have a northwesterly trend, parallel to that of the pre-Cambrian quartzites.

The foliation of the hornblende gneiss south of Quartz Creek has a north to northwesterly strike and dips steeply in either direction. North of Quartz Creek the strike is between north-northwest and north-northeast, except where the foliation parallels the contact of the quartz monzonite.

The hornblende gneiss and tonalite have the widest distribution of any rock type and occupy the central part of the Quartz Creek district.

These mafic rocks extend for a considerable distance to the northeast beyond the area mapped, where they have been described by Crawford _/. They

are the host rocks for a very large number of pegmatites, and numerous fine-grained granite dikes.

The hornblende gneiss and tonalite range from fine-grained to coarse-grained; the maximum grain size is about 0.20 inch. Textures or structures commonly found are: (1) prominent, well banded gneissic structure, (2) diabasic texture, (3) porphyritic texture, and (4) equigranular texture. Exposures of this rock are in general poor, and even where well exposed the textural changes are so great that in most places separation into mappable units was not feasible. Both rocks are dark gray to greenish black where fresh, and weather greenish gray to reddish brown. The hornblende content ranges from 20 to 80 percent, but most of the rock contains 50 to 75 percent hornblende. Some facies are unusually rich in hornblende and the rock may then grade into a hornblendite or perknite.

The minerals in the hornblende gneiss and in the tonalite are essentially the same, but the proportions of each vary widely. Hornblende, biotite, and feldspar are the only minerals that can be identified megascopical-

_/ Crawford, R. D., and Worcester, P. G., Geology and ore deposits of the Gold Brick district, Colorado: Colorado Geol. Survey Bull., vol. 10, pp. 27-28, 1916.

ly. In places much of the hornblende has altered to biotite. Andesine is the dominant light-colored constituent, but quartz and microcline are present locally. The accessory minerals are apatite, zircon, sphene, magnetite, epidote, chlorite, and sericite.

Much of the hornblende is in distinct dark-green euhedral crystals, but part is in frayed, ragged, pale-green grains that are altered largely to biotite. In one place it is altered to chlorite. Biotite is not found in some areas, but in others it is abundant. It forms as much as 60 percent of the rock, is commonly fresh, and occurs in brown prismatic crystals. Andesine (An₃₀-An₄₄) is poorly twinned and commonly is clouded with fine kaolin and sericite. The andesine crystallized after the hornblende in most places and fills the spaces between the hornblende crystals; in a few places the reverse is true. Quarts is present in most specimens constituting a maximum of 7 percent of the rock and occurs as small, clear grains with sutured borders. This mineral is interstitial to and cuts andesine; rarely they are micrographically intergrown. Because of the presence of a small amount of quartz, the rock is called a tonalite rather than a diorite, as used by Crawford _/. Microcline is present in a few places, but in most of

_ Crawford, R. D., and Worcester, P. G., op. cit., pp. 27-28, 1916.

the rock examined, it was absent. A trace to several percent of apatite and zircon are almost universally present as well-formed euhedral crystals associated with biotite. Epidote and sphene are found locally, usually where the hornblende is pale green and has been considerably altered. Magnetite occurs in irregular grains and is not common. Augite was noted in one specimen.

Quartz monzonite

Quartz monzonite crops out along the northern boundary of the Quartz Creek district. Pegmatites, similar in size and shape to those in horn-blende gneiss, are common near the outer edge of the intrusive. Farther into the mass the pegmatites are only a few inches thick and are very irregular in shape.

The quartz monzonite is a light- to dark-gray porphyritic rock that ranges in composition from quartz monzonite to granodiorite. Poor exposures make it difficult to map the variations of this rock in the field.

Mafic minerals (12 to 22 percent) are in clots and streaks composed of biotite, hornblende, zircon, sphene, magnetite, and apatite. Hornblende (0 to 15 percent), the dominant dark mineral, generally has been frayed and altered to biotite. Biotite (7 to 14 percent) occurs in small brown unaltered flakes and in clots or aggregates that appear megascopically to be large crystals. Apatite and magnetite (1 to 2 percent) commonly occur with biotite. Zircon, in trace quantities, is universally present as small crystals. Wedge-shaped brown sphene crystals locally make up as much as 5 percent of the rock. The leucocratic minerals are quartz, andesine, and microcline. Both andesine and microcline are in large phenocrysts and in fine grains in the groundmass. The feldspar content ranges from about 12 to 45 percent microcline and from about 30 to 65 percent andesine. The plagicclase has a composition of An31-An35. The microcline shows crosshatch twinning in most places. The andesine has some albite and pericline twinning and in some thin sections is covered by a thin film of kaolin and sericite. No quartz is observed megascopically, but in thin section small clear grains, interstitial to the feldspars, make up 4 to 15 percent of the rock. The quartz exhibits strain shadows and in many places has sutured

borders.

The quartz monzonite was intruded into the hornblende gneiss and in turn was cut by pegmatites. It is thus intermediate in age between pegmatite and hornblende gneiss. The age of the quartz monzonite in relation to the coarse-grained and fine-grained granite is not definitely known because the two rocks are not in contact. The following evidence, however, suggests that the quartz monzonite is older: (1) in many regions the differentiation of a batholith results in the early formation of mafic rocks. Subsequently, rocks of intermediate and granitic composition are formed as in the Quartz Creek district where the quartz monzonite is intermediate in composition between the granite and the earlier tonalite; (2) the gneissic texture in the hornblende gneiss parallels the contact on the quartz monzonite (pl. II). This implies that the quartz monzonite was intruded during metamorphism; whereas the coarse-grained granite cuts across foliation in many places; (3) the pegmatites that cut the granite and quartz monzonite have a composition similar to the granite and appear to have been derived from it rather than the quartz monzonite.

Coarse-grained granite

A large band of coarse-grained granite (pl. II) trends north-northwest across the district from the northeastern corner of sec. 22 to the northern border of sec. 33. This granite forms the prominent mountains on the southeast side of Quartz Creek. Another band of massive granite crops out a mile to the west and extends about one mile north of the southern boundary of the area mapped. These two granite masses converge several miles south of the Quartz Creek pegmatite district to form a large V. In addition to these two large granite bodies, numerous small bodies with exposed areas of from a few feet square to 2,000 feet by 800 feet, are scattered throughout

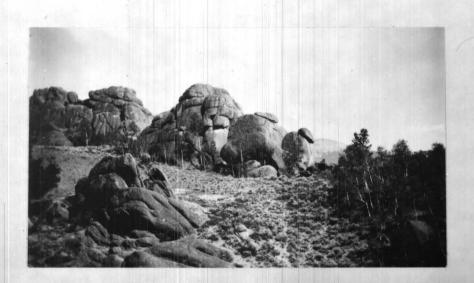
the hornblende gneiss and tonalite terrane. The area of most abundant small scattered granite intrusives is west of the main granite mass and trends north-northwest.

The granite is a pink porphyritic rock (pl. IV, A and B) that forms well-rounded outcrops. The phenocrysts are grains of pink microcline, 0.50 to 0.75 of an inch long, and grains of clear quartz, 0.25 to 0.50 of an inch long. In thin section the microcline phenocrysts show many small included crystals of diversely oriented microcline, quartz, biotite, and albite. The quartz phenocrysts are composed of several grains, commonly with sutured borders. The coarse-grained groundmass consists of interlocking grains of microcline, quartz, biotite, albite, magnetite, apatite, and zircon. Some specimens also contain sericite, epidote, and chlorite. The average composition of this rock is estimated to be microcline (71 percent), quartz (20 percent), biotite (8 percent), albite (1 percent), and less than 1 percent of magnetite, apatite, and zircon, and trace quantities of epidote, sericite, and chlorite. Apatite and zircon are most abundant in the biotite. Epidote commonly occurs near the biotite and the chlorite is derived from the biotite.

There are two less common varieties of the coarse-grained granite.

Granite gneiss occurs in a few isolated masses near the western edge of the district and is characterized by the parallel arrangement of elongate quartz and biotite crystals, granulation, and recrystallization of the quartz. Much of the biotite has been altered and only small wisps and discolored areas remain. The gneiss is believed to be a normal granite that has been metamorphosed by shearing.

A red granite occurs in small patches within the two main granite masses. It has no large phenocrysts, has many small vugs, contains only a



A. Coarse-grained granite along the divide between Wood Gulch and Quartz Creek.



B. Coarse-grained granite with two sets of joints at right angles.

few percent of quartz and a trace of biotite, and is high in albite. This rock is thought to be an albitized variety of the normal granite.

The granite is younger than the hornblende gneiss, tonalite, and preCambrian quartzites because it cuts these rocks or contains partly assimilated inclusions of them. On Indian Head, a large granite mass jutting
out into the valley of Quartz Creek (pl. II, sec. 4), numerous partly
assimilated fragments have large microcline porphyroblasts, and one piece
has well-rounded pebbles of conglomerate. The granite is in turn cut by
dikes of fine-grained granite and by pegmatite. Its relation to the quartz
monzonite is not clear, but the quartz monzonite is probably older than the
granite.

Fine-grained granite dikes and pegmatites

Pegmatite and fine-grained granite (pl. II) are found together in many places and cut both hornblende gneiss and the coarse-grained granite. The fine-grained granite dikes are cut by the pegmatites. The pegmatites are found throughout the district, except in the central parts of the two main granite masses. The fine-grained granite is much more restricted in distribution, and occurs in dikes in a north-northwesterly trending zone west of the largest granite intrusion. The same zone contains many small intrusive bodies of coarse-grained granite. A few fine-grained granite dikes are found in hornblende gneiss 200 feet from the northeast edge of the largest granite body.

The structure in the earlier pre-Cambrian rocks is followed in part by the granite dikes on the north side of Quartz Creek, which have a general trend of from north to N. 20° W. (pl. II). South of Quartz Creek the dikes have a general northeasterly trend, but range from N. 20° W. to N. 50° E.

The pegmatites form both long narrow dike-like bodies and irregular masses. The dikes trend, except in sec. 33, from N. 15°-60° E., cutting across the earlier structure. In the vicinity of the southeast corner of sec. 33, T. 50 N., R. 3 E. the pegmatites have an average trend of N. 35° W. and are described in detail in a succeeding part of this report.

The fine-grained granite is pink and has a grain size of about 0.015 of an inch. The dikes range in width from a few inches to 180 feet and in length from a few feet to 2,700 feet. The contacts with the surrounding rock are sharp, and the granite forms prominent outcrops. The rock is made up almost entirely of leucocratic microcline, quartz, and plagioclase. Microcline (20 to 60 percent) has crosshatch twinning. Clear quartz (15 to 40 percent) forms irregular grains, many with sutured borders, and is interstitial to plagioclase and microcline. The plagioclase (20 to 40 percent) is albite (An_t) and occurs in crystals coated with kaolin and as inclusions in microcline crystals. Biotite is the dominant dark mineral ranging from a trace to about 5 percent; the average is less than 1 percent. Ragged grains of muscovite, commonly on feldspar, is as much as a few percent in some places. A few euhedral crystals of apatite and irregular-shaped grains of magnetite are present in some specimens.

A gray facies of this rock in secs. 8 and 17 differs from the average fine-grained granite in that it contains about 5 percent biotite and has andesine (An₃₂-An₃₄) feldspar. It is poorly exposed and its relation to the other rocks is not known.

The fine-grained granite is related in age to the coarse-grained granite, and probably was derived from the same magma, but at a later date. This age relationship is indicated by their areal distribution. The small coarse-grained granite bodies and the fine-grained granite dikes crop out in the same north-northwesterly trending band west of the main granite mass and

the fine-grained granite dikes also occur in a narrow zone along the northern contact of the largest granite mass. Both rocks are of the same mineral composition, but the fine-grained granite is commonly richer in plagioclase and poorer in microcline and may represent a more sodic fraction of the magma.

Mesozoic rocks

Morrison formation

The Morrison formation unconformably overlies the pre-Cambrian and is conformably overlain by the Dakota sandstone along the western and eastern edges of the Quartz Creek district. The Morrison formation is covered in more than 90 percent of the area and the outcrops are commonly of the more resistant sandstone lenses.

This formation is composed of a basal and an upper sandstone that are separated by vari-colored shale. The basal sandstone rests on the pre-Cambrian and closely resembles the Dakota sandstone in appearance. It is white to tan and weathers buff to yellowish-brown. The quartz grains are subrounded, and a few beds are quartzitic. The middle unit of the formation rarely is exposed. It is composed of green, brown, and reddish shales with a few thin limestone and sandstone beds. Above the shales is a white fine-grained sandstone flecked with iron oxide. This rock is prominently cross-bedded, is usually friable, and the individual quartz grains are well rounded. This sandstone is conformable with the basal pebble conglomerate of the Dakota sandstone.

At no place in the area is a complete section of the Morrison formation exposed, but a thickness of 355 feet was measured along the west side of Alder Creek, in sec. 36, T. 50 N., R. 2 E., from the top of the underlying

pre-Cambrian (as determined by float) to the base of the Dakota formation.

The thickness of this formation also was measured by Dings _/ in the south-

___ Dings, McClelland, Personal communication, 1949.

western corner of the adjacent Garfield quadrangle. His measurements, made under equally difficult conditions, with the exact position of the upper and lower limits inferred, indicate the thickness of the Morrison to be between 315 and 375 feet.

No fossils were found and the identification of the Morrison formation in the Quartz Creek district is based on its lithologic similarity to this formation in other areas.

Dakota sandstone

The Dakota sandstone is well exposed in a series of cliffs that border Alder Creek (pl. II). This nearly flatlying formation crops out in a series of prominent steplike cliffs. In plate V, A, it is shown capping the Morrison formation along the western border of the district; it also crops out east of the mapped area.

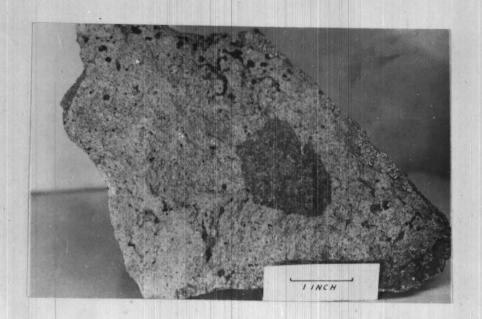
The Dakota is composed of a basal pebble conglomerate and an upper sandstone. The conglomerate is composed of subrounded to rounded pebbles averaging 0.25 of an inch in diameter. The pebbles are for the most part quartz, with subordinate amounts of black chert and red jasper. In part the conglomerate is arkosic though much of the feldspar has altered to clay. The upper part of this unit is quite friable and commonly cross-bedded, whereas the lower part locally is cemented with chalcedony and is very resistant.

The upper unit of the formation is composed almost entirely of sandstone, but in its upper part contains thin beds of fine-grained black to

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Plate V



A. Cliff of Dakota sandstone.



B. Tuff with large volcanic fragment.

gray fissile shale a few inches to 1.5 feet thick. This sandstone is composed dominantly of subrounded grains of quartz, and subordinately of orthoclase. The rock ranges from a true arkose with about 25 percent feldspar to an almost pure quartz sandstone. The cliff-forming units are compact sandstone, but much of the unit is soft and friable. Locally it has been indurated to quartzite. The sandstone is white to gray and weathers buff or yellowish brown. Commonly it contains small spots of limonite, and one bed in the upper part contains radiating 1-inch spheroids of limonite pseudomorphous after pyrite.

The Dakota sandstone is not completely exposed in the area mapped and the upper surface has been eroded. The maximum thickness obtained from the six sections measured is 183 feet; the basal conglomerate is 33 feet thick.

The only fossils found in this formation were a poorly preserved unidentifiable gastropod and a few fragmentary casts of plant stems.

Tertiary rocks

Tuff

A white tuff is well exposed on the south side of Quartz Creek, in the southeast corner of the Quartz Creek district (sec. 8), where it forms a small cliff overlying the Dakota sandstone. At two other smaller areas of outcrop, in sec. 16 on the south side of Wood Gulch and in sec. 5 on the north side of Quartz Creek, the tuff overlies pre-Cambrian rocks.

The tuff is a porous, white flaggy rock occurring in layers 1 to 2 inches thick. The layering dips from 4 to 23 degrees. A few subrounded fragments of darker volcanic rock (pl. V, B) are enclosed in an aphanitic matrix containing phenocrysts, one-thirty second of an inch long, of plagioclase and biotite. The orientation of the biotite in general is parallel to the layering.

In thin section, the tuff has a clastic texture and shows many scattered phenocrysts in a brown cryptocrystalline groundmass. The phenocrysts comprise about 35 percent of the rock. Andesine-labradorite (more than 30 percent) commonly occurring as angular fragments or rarely as euhedral crystals, is the most common phenocryst. Some of the plagioclase crystals show zoning. Long fibrous crystals of brown biotite (3 percent) is the chief dark mineral. Next in abundance is black anhedral magnetite (1 percent). Other minerals in the approximate order of their abundance are: hornblende, quartz, scapolite, sphene, apatite, and zircon.

The tuff outcrops are erosional remnants, a few tens of feet thick, of a thicker and more extensive tuff bed. The tuff that overlies the Dakota sandstone along the east side of Quartz Creek has a minimum thickness of 83 feet. The tuff is younger than the faulting which brought pre-Cambrian rocks against the Dakota sandstone, because the tuff is deposited across the fault line with no apparent displacement. It is overlain along Wood Gulch by glacial till, presumably of Pleistocene age. The tuff, therefore, has been tentatively designated as Tertiary.

Pleistocene and Recent deposits Glacial till

Pleistocene (?) glacial till overlies the other formations along both sides of Quartz Creek and Wood Gulch, and fills the broad valley of Quartz Creek where it is covered by a foot or less of soil. On the south side of Quartz Creek the till is quite thin and pegmatite outcrops protrude through it. There is considerable difference in the altitude at which the till was deposited. The highest position is on the north side of Quartz Creek at 8,700 feet; on the south side the altitude is 8,250 feet. The till deposits on the north side are part of the lateral moraine whereas those on the south

side are till ridges in the valley moraine. The till near the mouth of Wood Gulch appears as thin irregular patches which seem to be remnants of a broad valley moraine.

The till is composed of clay, fine sand, pebbles, and boulders as much as 3 feet in diameter. The boulders are a heterogeneous mixture of several rock types and differ from place to place. On the south side of Quartz Creek and along Wood Gulch, hornblende gneiss and tonalite are the dominant rock types in the till and in places form more than 80 percent of it. Near the mouth of Alder Creek on the north side of Quartz Creek the till consists of pegmatite (30 percent), Dakota and Morrison sandstone (30 percent), fine-grained granite (20 percent), rhyolite porphyry (10 percent), and hornblende gneiss (10 percent). Other identifiable boulders include chert, pre-Cambrian quartzite, Sawatch quartzite, epidote rock, quartz monzonite, basalt, massive quartz, and andesite.

Alluvium

Alluvium forms a narrow strip in the bottom of most of the valleys in the Quartz Creek district. Along Quartz Creek this strip is 1/8 to 3/4 of a mile wide, and extends northeastward across the entire district. The alluvium is dominantly fine silt, 4 to 8 inches thick, and overlies glacial deposits along most of Quartz Creek.

STRUCTURAL GEOLOGY

The structure of the older pre-Cambrian rocks of the Quartz Creek district has a general northwest trend, which is cut by stocks and batholiths of younger pre-Cambrian granite and quartz monzonite. Mesozoic and later rocks are flatlying and are cut by several faults in the southern part of the district.

The general trend of the pre-Cambrian metasedimentary rocks is north-west, with a steep dip southwest. The foliation of the pre-Cambrian horn-blende gneiss strikes northwest to northeast and dips steeply. On the southeast side of Quartz Creek and along the western edge of the district the foliation trends northeast and dips from 70° SE. to 59° NW. Around the edge of the quartz monzonite intrusion the foliation parallels the contact and dips steeply away from it. In the northern part of the district, the foliation strikes northeast and dips from steeply southwest to vertical.

The large granite mass dips steeply to the northeast along its northeastern side. On the west, however, the contact was not exposed, but due to the innumerable small stocks along this side (pl. II) it is believed that the granite underlies the schist at shallow depth. The contact of the quartz monzonite was not exposed but the strike of the foliation of the hornblende gneiss is oriented parallel to that of the contact, and it is probable that the dip is also parallel.

Most of the pegmatites trend northeast along joints and cut across the foliation of the older rocks. Groups of parallel lenticular pegmatites with this trend are common (pl. VI, A).

Faults are difficult to recognize in the pre-Cambrian rocks except where pegmatites have been cut and offset. The displacement observed ranges from a few inches to 4 feet. Drag folds and local disruptions in the foliation also may have been the result of unrecognized faulting.

Two sheared and mineralized fractures were mapped in the hornblende gneiss. The larger of these is south of Quartz Creek, 250 feet east of the Buckhorn pegmatite (No. 659). The second shear zone is in the northwestern part of the area mapped, where the southern part of pegmatite No. 1199 has been displaced about 3 feet to the west.



A. Pegmatites showing regional trend. Brown Derby mine in the background.



B. Anticline in center of picture is a down faulted block of Dakota sandstone.

In the Mesozoic sediments faults are more readily recognized. A major fault separates Dakota sandstone from hornblende gneiss in the southwestern part of the district (pl. II) and trends N. 20°-42° W. A vertical displacement of 410 feet was measured on the west side of Alder Creek, the southern block having moved downward with respect to the northern block. On the west side of Alder Creek the Dakota sandstone has been sharply upturned by drag of the beds at the fault. In the southwestern part of the district, along State Highway 162, a small segment of Dakota sandstone has been downfaulted between the large fault and two smaller ones to the level of the highway and folded into a gentle anticline (pl. VI, B).

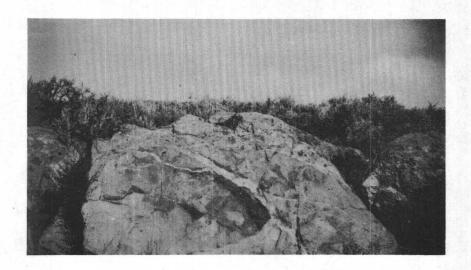
PEGMATITES

Size and shape of pegmatites

The pegmatites of the Quartz Creek district range in size from bodies a few inches wide and a few feet long to bodies like the Black Wonder pegmatite, 12,600 feet long by 6,700 feet in maximum width. Most pegmatites range from 100 to 400 feet in length, but 37 bodies are over 1,000 feet long. The two largest pegmatites are the Bucky deposit, 4,000 feet long by 2,600 feet in maximum width, and the Black Wonder, both of which are very irregular and have many small branches. The small pegmatites are commoner in the granite and quartz monzonite.

The pegmatites in the Quartz Creek district can be classified, on the basis of shape, as: (1) lenticular, (2) lenticular-branching, (3) oval, and (4) irregular. Examples of each are shown in figures 2 to 6 and plate VII, A and B, including both the extreme variations and the average shape in each type. Each of these examples represents many more pegmatites of similar shape. The lenticular pegmatites are 2.3 times more common than ir-

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A. Small branching pegmatite cutting fine-grained granite.



B. Large branching pegmatite (No. 250) cutting horn-blende gneiss.

and the same

regular pegmatites—the second most abundant type. The general order of frequency is one oval pegmatite to 2.3 lenticular-branching pegmatites, 218 irregular pegmatites, and 6.6 lenticular pegmatites.

Comparison of pegmatite shapes in this district with shapes in the Black Hills _/ and other pegmatite districts has shown that the shape of a

_/ Page, L. R., et al., Pegmatite investigations 1942-1945, Black Hills, South Dakota: U. S. Geol. Survey Prof. Paper (In preparation).

granitic pegmatite is controlled primarily by the type and competency of country rock, and as the amount of material intruded becomes large it controls the shape of the pegmatite to a greater degree.

The pegmatites in the Quartz Creek district for the most part are intruded into granite, quartz monzonite, hornblende gneiss, and tonalite. The hornblende gneiss and tonalite have a similar composition, but the hornblende gneiss is foliated and the tonalite is equigranular. Both rocks are competent and the pegmatites tend to follow fractures and joints that cut the poor to well-developed foliation in the hornblende gneiss. Though the pegmatites intruded into hornblende gneiss are usually well exposed, the adjacent gneiss is rarely seen. Wherever the foliation of the hornblende gneiss was exposed adjacent to the pegmatite the angle between the foliation plane and the side of the pegmatite was measured; the results are plotted in figure 8 on a bar graph. This graph indicates that there is no constant angle at which the pegmatites cut the foliation of the country rock; though it is most common up to 60 degrees. The irregularity of the pegmatites and their numerous changes in direction point to the emplacement of the pegmatites along irregular fractures and joints. The hornblende gneiss and tonalite are too poorly exposed to allow measurement of any overall joint systems. The largest body of coarse-grained granite is

well exposed and 639 joints were measured (fig. 9). In local areas of several hundred square feet where 50 to 60 joints are exposed, they are related to 2 or 3 well-developed sets of joints. Over the entire granite body, however, 639 joints show a random orientation. The main granite body is cut by pegmatites only in its northwestern end, where the lenticular, and lenticular-branching pegmatites trend N. 45° W. In the hornblende gneiss and tonalite the pegmatites trend from north to N. 45° E. (pl. II). The trend of the lenticular pegmatites in the mafic rocks is quite uniform over the whole district. This points to a district-wide joint system in the hornblende gneiss and tonalite, whereas the joint systems in the granite vary from one locality to the next. A probable explanation of this peculiar feature is that the joint system in the mafic rocks antedates the intrusion of the granite, and that the jointing in the granite was developed by local stresses at the time of the intrusion.

In comparing the various types of country rock to the shapes of pegmatites it was found that in areas of fractured competent rocks, lenticularbranching pegmatites predominate and are found along intersecting fractures; to a lesser extent irregular pegmatites and thin sinuous bodies
characterize such terranes. The composition of the country rock has little
to do with the shape of the pegmatite, provided the rocks being compared
are of equal competency. Table 1 shows the frequency of occurrence of each
shape compared to the oval shape in each of the three most common types of
country rock, hornblende gneiss and tonalite, coarse-grained granite, and
quartz monzonite.

Table 1,-Ratio of pegmatite shapes to the oval type in different kinds of country rock.

The angular record reco	Pegmatité shapes										
Rock type	Lenticular	Lenticular-branching	Irregular	Oval							
Hornblende gneiss and tonalite	5.8	2.6	2,4	1.0							
Coarse-grained granite	6.2	2.6	3.8	1,0							
Quartz monzonite	6.4	2.0	6.0	1.0							

All pegmatites that cut more than one rock type are omitted (fig. 7). Except for the larger number of irregular pegmatites in the quartz monzonite the ratios are remarkably similar. The higher ratio of irregular pegmatites in the quartz monzonite probably can be correlated there with the greater number of large pegmatites in this area. This is discussed in more detail in a succeeding paragraph. The country rocks differ greatly in their mineralogy texture, and chemical composition, and yet the shapes of the pegmatites show little variation. The rocks of this district have one important characteristic in common: they are all tight, brittle, and competent. The effect on the shape of the pegmatites where the host rock possesses even minor foliation is quite striking. Table 2 shows the frequency of occurrence of the different shapes as a ratio related to the oval shape in hornblende gneiss and tonalite.

Table 2.—Ratio of various pegmatite shapes to the oval type in hornblende gneiss and tonalite.

	Pegmatite shapes									
Rock type	Lenticular	Lenticular-branching	Irregular	Oval						
Hornblende gneiss	8.7	2,0	2,8	1.0						
Tonalite	12.0	9.2	2.2	1.0						

Foliation has a profound effect in simplifying the shapes of the pegmatites

by decreasing the number of branching types; the lenticular-branching type is 4.6 times more common in the tonalite than in the hornblende gneiss.

Pegmatites in similarly competent rocks are found also in hornblende schist in the Bridger Mountains _/, Wyoming, and in Pikes Peak granite in the

_/ McLaughlin, T. G., Pegmatite dikes of the Bridger Mountains, Wyoming: Am. Mineralogist, vol. 25, pp. 46-60, 1940.

Eight Mile Park, Colo._/

_____ Heinrich, E. W., Pegmatites of Eight Mile Park, Fremont county, Colorado: Am. Mineralogist, vol. 33, pp. 420-448, 1948.

Pegmatites in incompetent rocks such as mica schist are in general concordant with the foliation and were intruded by shouldering apart the country rock. Pegmatites of this type are most commonly lenticular, but other common forms are troughlike, arcuate, and teardrop. Lenticular-branching pegmatites are extremely rare. The schistosity of the wall rock is commonly conformable around the entire pegmatite.

The second factor influencing the shape of pegmatites is in the quantity of material intruded. With the intrusions of large quantities of pegmatite material the control of the structures in the country rock on the shape is usually obscured and the body becomes an irregular stocklike mass. The directional control that fractures had on large pegmatites like the Black Wonder, Bucky, or Buckhorn was obliterated by the large quantity of material and is found only in the small stringers that extend outward from the main mass.

Whether pegmatites are in competent or incompetent country rocks makes a great difference when it comes to predicting their position, shape, and attitude underground and in calculating ore reserves. Those in incompetent rocks may be projected with some confidence, whereas, pegmatites in com-

petent rocks such as in the Quartz Creek district, can be predicted only if the attitudes of the controlling fractures are known.

Internal structure

The recognition of distinct lithologic and structural units within pegmatites dates back for many years. Hunt _/, who noted a remarkable

_/ Hunt, T. S., Notes on granitic rocks: Am. Jour. Sci., 3d ser., vol. 1, pp. 89, 182-186, 1871.

banded arrangement "formed by successive deposits of mineral matter" at Brunswick, Topsham, and Newry, Maine, appears to be the first American geologist to recognize a systematic internal structure in pegmatites.

Many early authors of 25 to 40 years ago referred to segregations, veins, layers, bands, and streaks. An excellent historical review of these early writings is given by Cameron, Jahns, McNair, and Page _/. Until about 1940

_/ Cameron, E. N., Jahns, R. H., McNair, A. H., and Page, L. R., Internal structure of granitic pegmatites: Econ. Geology Mon. 2, pp. 10-13, 1949.

_/ See for example:

Fraser, H. J., Paragenesis of the Newry pegmatite, Maine: Am. Mineralogist, vol. 15, pp. 349-364, 1930.

Hess, F. L., Pegmatites: Econ. Geology, vol. 28, pp. 447-462, 1933.

The natural history of pegmatites: Eng. and Min. Jour.Press, vol. 120, pp. 289-298, 1925.

Landes, K. K., Paragenesis of the granitic pegmatites of central Maine: Am. Mineralogist, vol. 10, pp. 355-411, 1925.

[,] Sequence of mineralization in Keystone, South Dakota pegmatites: Am. Mineralogist, vol. 13, pp. 519-530, 537-558, 1923.

, Criteria of age relations of minerals: Econ. Geology, vol. 27, p. 211, 1932.

Landes, K. K., Origin and classification of pegmatites: Am. Mineralogist, vol. 18, pp. 33-56, 95-103, 1933.

, Colorado pegmatites: Am. Mineralogist, vol. 20, pp. 319-333, 1935.

Schaller, W. T., The genesis of lithium pegmatites: Am. Jour. Sci., 5th ser., vol. 10, pp. 269-279, 1925.

, Mineral replacement in pegmatites: Am. Mineralogist, vol. 12, pp. 59-63, 1927.

, Pegmatites: Ore deposits of the western states, pp. 144-151, Am. Inst. Min. Met. Eng., New York, 1933.

and put little emphasis on their structure.

After 1940, because of the wartime need for pegmatite minerals, the U. S. Geological Survey made numerous studies of the internal mineralogic and structural units in pegmatites. As the economical concentrations of valuable minerals in pegmatites tend to be concentrated in rock units distinct from the adjacent barren units, detailed mapping and interpretation of various pegmatite units have proved of much aid in exploration, development work, and mining _/. Drilling on the basis of structural interpreta-

[/] For example see:

Smith, W. C., and Page, L. R., Tin-bearing pegmatites of the Tinton district, Lawrence County, South Dakota: U. S. Geol. Survey Bull. 922, pp. 595-630, 1941.

Olson, J. C., Mica-bearing pegmatites of New Hampshire: U. S. Geol. Survey Bull. 931-P, pp. 363-403, 1942.

Bannerman, H. M., Structural and economic features of some New Hampshire pegmatites: New Hampshire Mineral Resources Survey, pl. 7, New Hampshire State Planning and Development Commission, Concord, pp. 1-22, 1943.

Cameron, E. N., Larrabee, D. M., McNair, A. H., Page, J. J., and Shainin, V. E., Structure and economic characteristics of New England mica deposits: Econ. Geology, vol. 40, pp. 369-393, 1945.

Johnston, W. D., Jr., Beryl-tantalite pegmatites of northeastern Brazil: Geol. Soc. America Bull., vol. 56, pp. 1015-1070, 1945.

Jahns, R. H., Mica deposits of the Petaca district, Rio Arriba County, New Mexico: New Mexico Bur. Mines, Bull. 25, 293 pp., 1946.

Cameron, E. N., Jahns, R. H., McNair, A. H., and Page, L. R., Internal structure of granitic pegmatites: Econ. Geology Mon. 2, 115 pp., 1949.

Hanley, J. B., Heinrich, E. W., and Page, L. R., Pegmatite investigations in Colorado, Wyoming, and Utah: U. S. Geol. Survey Prof. Paper 227, 1950.

tion of the internal units has given excellent results _/.

The internal units of pegmatites have been classified as (1) zones,

(2) fracture fillings, and (3) replacement bodies _/. Many of the pegmatites

_/ Cameron, E. N., Larrabee, D. M., McNair, A. M., and Stewart, G. W., Characteristics of some New England mica-bearing pegmatites (abstr.): Econ. Geology, vol. 39, p. 89, 1944.

Jahns, R. H., Mica deposits of the Petaca district, Rio Arriba County, New Mexico: New Mexico Bur. Mines Bull. 25, pp. 39-51, 1946.
Heinrich, E. W., Pegmatites of Eight Mile Park, Fremont County,

Heinrich, E. W., Pegmatites of Eight Mile Park, Fremont County, Colorado: Am. Mineralogist, vol. 33, pp. 436-442, 1948.

Cameron, E. N., Jahns, R. H., McNair, A. H., and Page, L. R., Internal structure of granitic pegmatites: Econ. Geology Mon. 2, pp. 13-97, 1949.

of the Quartz Creek district differ from those in other pegmatite areas in that in addition to these three units they may contain the primary internal structure designated as banding in this paper.

Zones

The zones of a pegmatite in ideal development are concentric shells about an innermost zone or core; in many places they are incomplete, however, forming only along one end or in one part of the pegmatite. Zonal structure is formed during the primary consolidation of the pegmatite magma and may be cut by fracture fillings and replacement bodies. Zones have been classified _/ as: (1) border zones, (2) wall zones, (3) intermediate zones,

_/ Jahns, R. H., op. cit., p. 42, 1946.
Cameron, E. N., Jahns, R. H., McNair, A. H., and Page, L. R., op. cit., p. 20, 1949.

and (4) cores.

Border zones are fine-grained selvages that in most pegmatites are a few inches or less in thickness. Most are of little significance in the mining or quarrying of pegmatites, and hence in the industry are not distinguished from the adjoining wall zones that are coarser grained and much thicker. Although they actually are the second zones from the margins of pegmatite bodies, they are designated as wall zones in recognition of terminology firmly established in the pegmatite mining industry. The innermost zone or core occurs at or near the center of the pegmatite either as an elongate lens or a series of disconnected segments. Any zone between the core and the wall zone is an intermediate zone. Any number of intermediate zones can exist, but few pegmatites contain more than three. If the core is not exposed at any one level, the innermost exposed zone may be identified erroneously as a core.

Banding

Banding is the name given to the layered structures forming pegmatite units that differ in mineralogy, texture, or both and tend to have a non-concentric arrangement within pegmatite bodies. Banding in a pegmatite may divide the body either across or along the strike. Several distinct types of banding are recognized in the Quartz Creek district.

Banding parallel to strike.—Pegmatites in which banding is parallel to the strike and dip of the body are called layered pegmatites (fig. 11). The distinct bands or layers are mappable units of definite mineralogy or texture and are not repeated. The layered pegmatites commonly consist of several tabular units whose contacts are approximately parallel to the hanging-wall and footwall sides of the pegmatite. These layers differ from zones in that there is no repetition of units on the other side of

the pegmatite. Pegmatites composed of two layers are by far the most common type in this district. These units commonly extend the entire length of the pegmatite and are from 1 to 30 feet thick. This type of banding is confined to narrow lenticular and lenticular-branching pegmatites or to a narrow lenticular part of irregular pegmatites. It is not found in thick parts of irregular pegmatites or oval pegmatites. The distinct upper and lower units in many of these layered pegmatites can be distinguished in only certain parts of the body and merge along strike into a single unit. Where two layers merge, or telescope, the unit formed has the bulk composition of the two combined layers and a texture intermediate between that of the upper and lower layers. In pegmatite No. 548, for example, an upper layer, consisting of perthite (50 percent), quartz (33 percent), albite (15 percent), and muscovite (2 percent), and a lower layer, consisting of albite (74 percent), quartz (20 percent), perthite (3 percent), and muscovite (3 percent), become progressively less distinct to the south and finally merge into a single unit, consisting of albite (63 percent), quartz, (20 percent), perthite (15 percent), and muscovite (2 percent).

Banding across strike.—Many pegmatites are banded across the strike into two or more mappable units differing in mineralogy, texture, or both (fig. 12). These are designated as pegmatites showing variation in composition along strike. Banding across strike results in two or more pegmatite units that have their contact at an angle to, rather than parallel to, the strike of the pegmatite units. In such pegmatites the bands occupy the full width of the body, and are from 20 to several hundred feet across parallel to the strike of the pegmatite body. The units have the shape of whatever part of the pegmatite they occupy; thus, one unit may occupy the short, thin lenticular part and another long, irregular, bulbous part of a

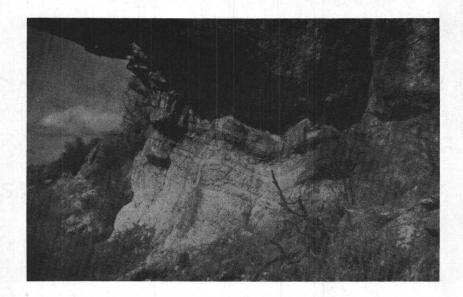
pegmatite. Banding across strike has been found only in lenticular and lenticular-branching pegmatites and all such bodies contain either two or three bands (fig. 12).

Multiple banding.—Some pegmatites in the Quartz Creek district are composed of innumerable very thin bands that differ in texture, mineralogy, or both. The bands are rarely mappable on ordinary scales. This type of banded rock has been described as "line rock" in the Pala district, Calif.

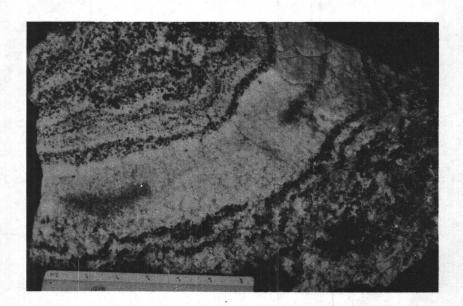
____ Schaller, W. T., Genesis of lithium pegmatites: Am. Jour. Sci., 5th ser., vol. 10, p. 273, 1925.

Line rock in the Quartz Creek district is characterized by the repetition of bands of minerals from 0.01 inch to 0.4 foot thick; the average thickness is less than 0.5 inch. The banding in most places, as in pegmatite No. 670, is parallel to the strike of the body, but in a few places it cuts across the strike. Line rock is found commonly as patches in a small part of the pegmatite. No pegmatite with the exception of pegmatite No. 670 contains more than 15 percent of line rock, and most of them contain less than 1 percent. Line-rock-bearing pegmatites, therefore, are not classified separately.

Line rock is most common in albite-rich pegmatites where the most obvious banding is caused by layers of garnet and muscovite as much as one-quarter of an inch thick (pl. IX, A and B). Layers of fine-grained albite-quartz pegmatite are interspersed with coarser layers of perthite-quartz-albite pegmatite (a half inch to about 4 inches thick). The layers of perthite-quartz-albite pegmatite are lenticular and usually pinch out within short distances. Other layers may occur above or below to form an echelon pattern. Rarely the albite-quartz pegmatite forms the lenticular units in line rock. The banding may end abruptly against large crystals or



A. Line rock in the lower part of pegmatite No. 670. Fine dark layers are small brown garnets.



B. Line rock from pegmatite No. 461. Black layers are garnet and muscovite and white layers are mainly albite with a little quartz.

an aggregate of minerals (fig. 14). Line rock is most common adjacent to the walls of the pegmatite, especially on the footwall side. Thin layers of garnet and albite were noted terminating abruptly against euhedral perthite crystals in several places. The perthite is veined by albite along fractures and was either entirely crystallized or at least partly crystallized before all the albite was deposited. Thus, the perthite is not a late mineral which cut off the layering, but rather a buttress against which the layering stopped. The arrangement of garnets and micas in bands suggests at least local movement of the pegmatite magma with rapid minor changes in composition so that first one mineral would be crystallizing and then another. Thus a row of garnets might be formed in an area of movement up to a projecting perthite crystal, which would either turn the current or deposit a row up to its side and another along its top. This row on top might be swept off its more exposed position where narrowing of the channel caused the current to be swifter, or it might be incorporated into the perthite upon further growth. The alternating lenses of perthitequartz-albite pegmatite in fine-grained albite-quartz pegmatite suggests zoned multiple pegmatites. They might have been formed by crystallization from lenses of trapped liquid.

Line rock is common in many pegmatites in other areas: the Crystal Mountain district, Colo._/, the Middletown district, Conn., the Pala district,

Calif._/, the Eight Mile Park district, Colo._/, and the Bridger Mountains,

_____ Schaller, W. T., op. cit, pp. 272-273.
____ Heinrich, E. W., Pegmatites of Eight Mile Park, Fremont County,
Colorado: Am. Mineralogist, vol. 33, p. 448, 1948.

_/ McLaughlin, T. G., Pegmatite dikes of the Bridger Mountains, Wyoming: Am. Mineralogist, vol. 25, pp. 62-63, 1940.

Fracture fillings

Fracture fillings are tabular bodies that extend from inner units into outer units of the pegmatite. In places they connect directly to the core.

Fracture filling units are common in pegmatites of the Quartz Creek district but are usually small; many are only a few feet in length. Most of these units are only a minor part of a pegmatite, though there may be several in a single pegmatite. Discontinuous core segments and fracture fillings are difficult to distinguish in some irregular pegmatites.

Most of the fracture fillings are coarse-grained and consist predomimently of perthite and quartz, or quartz pegmatite. In pegmatite No. 1096 a fracture filling of massive quartz extends from the core across the wall zone.

Replacement units

No mappable replacement units were found in the pegmatites of the Quartz Creek area, although there are several places where small areas were replaced along fractures. Replacement units form by the replacement of pre-existing consolidated pegmatite with later material. The interaction of two minerals or of a mineral with the rest solution during the process of crystallization is not considered as replacement in this paper. The embayment of one mineral by another and the filling of small fractures have been given as criteria of replacement, but these textures also can be formed if an early-formed mineral is corroded by the rest solution and

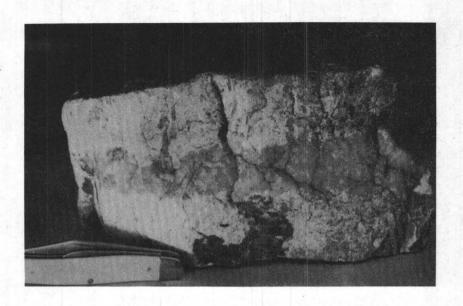
subsequently coated by a later mineral.

The criteria used to distinguish a replacement body are, therefore, the presence of relic textures or structures of the pre-existing rock that indicate essentially complete consolidation. Where a pegmatite is not zoned or where no pre-existing textures or structures remain it may be very difficult to recognize a replacement body.

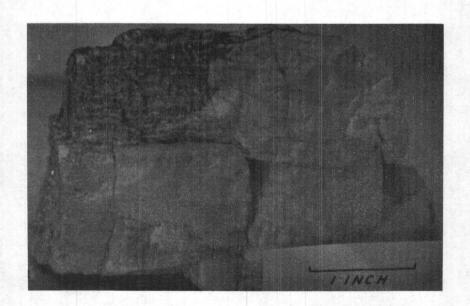
The interaction of one mineral on another during the crystallization process is more pronounced in pegmatites than in other igneous rocks because of the large grain size of the crystals which magnifies the embaying of one mineral by another and the long crystallization period caused by the presence of volatiles. Thus, a crystal of one mineral may be partly or completely grown before the equilibrium in the solution will permit a second mineral to start crystallizing. This crystal may form around the first crystal, grow out from it or, in the new equilibrium, the first mineral may be soluble and may be replaced by the second. Evidence of this sort does not prove or disprove the presence of a replacement unit. An excellent example of muscovite selectively replacing perthite and leaving narrow albite stringers of the perthite intergrowth unreplaced, is illustrated in plate VIII, B. The pegmatite was essentially homogeneous and shows no evidence of a separate replacement body.

Types of pegmatites

Pegmatites may be divided into homogeneous and heterogeneous pegmatites. The homogeneous pegmatites are simple aggregates of feldspar, quartz, and accessory minerals which cannot be divided into contrasting units on the basis of mineralogy and texture. These pegmatites form the great bulk of pegmatites in many regions, such as the Quartz Creek district, Colo., Black Hills region, S. Dak., Spruce Pine district, N. C._/, and the Crystal



A. A thin zoned pegmatite with an albite-quartz wall zone and a quartz core (grey).



B. Replacement of perthite by fine-grained muscovite (black).

Note the unreplaced albite lamillae of the perthitic structure in muscovite.

Mountain	district,	Colo	/ Hon	nogeneous	pegmatites	commonly	form	relative-
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ly small dikes and rarely contain minable concentrations of economic minerals. Because they lack the economic and rarer minerals, they have received in the past little consideration by mineralogists and geologists. Most of the pegmatite literature is devoted to descriptions of pegmatites containing the rarer minerals. Recently, however, the U. S. Geological Survey has undertaken regional mapping of pegmatite districts in South Dakota, other parts of Colorado, North Carolina, and Connecticut—similar to that described in this report—that should result in obtaining a broad knowledge of the character and distribution of all types of pegmatites. This work may yield the much-needed data required to understand the relationships of homogeneous and heterogeneous pegmatites.

Heterogeneous pegmatites are those that can be divided into different rock units on the basis of mineralogy, texture, or both.

Homogeneous pegmatites

Homogeneous or one-unit pegmatites form the great bulk of the pegmatites in the Quartz Creek district. Out of more than 1,800 pegmatites, 78 percent are homogeneous or "unzoned". Homogeneous pegmatites occur as lenticular, lenticular-branching, oval, and irregular bodies. Only a few of the larger irregular pegmatites are zoned; they contain one or more discontinuous cores. Examples of homogeneous pegmatites are shown in figures 2, 3, 4, 5, and 6. The dominant minerals are plagioclase, quartz, and

perthite. Most of the pegmatites contain as much as 2 percent of one or more of the following minerals: muscovite, garnet, biotite, and magnetite.

Beryl and tourmaline are present in some pegmatites but constitute only a small fraction of 1 percent.

Heterogeneous pegmatites

Zoned pegmatites.—Zoned pegmatites form roughly 14 percent of the pegmatites in the area (figs. 7 and 10); most have only a core, a wall zone, and a narrow border zone. Most of the border zones in the Quartz Creek district are only a fraction of an inch thick and are cooling selvages. Owing to their thinness they were mapped with the wall zones.

The mineralogy and texture of the wall zone usually resembles the homogeneous pegmatites in the immediate vicinity. The core, except where the wall zone is predominantly graphic granite, is commonly coarser grained, and contains more perthite, or quartz, or both, than the surrounding wall zone. Cores consisting only of massive quartz are very common (pl. VIII, A) in the northwestern part of the district. Cores of perthite-quartz pegmatite also are common. There are 14 pegmatites consisting of a wall zone and a core, that have cores of cleavelandite-quartz, or cleavelandite-lepidolite-quartz pegmatite. These pegmatites include the Bazooka, White Spar No. 2, and the Brown Derby No. 2 and No. 3 pegmatites, that have been described previously by Hanley, Heinrich, and Page J. Not all the

lepidolite-bearing and cleavelandite-bearing pegmatites are zoned pegmatites; these minerals are also found in small homogeneous pegmatites,

layered pegmatites, and pegmatites which vary in composition along their length. The cores in small pegmatites may form a large proportion of the pegmatite (No. 267, fig. 10), but in large pegmatites they usually make up 1 percent or less of the total rock (fig. 7). Pegmatites having an intermediate zone as well as a core and a wall zone are rare. Only 7 pegmatites in the Quartz Creek district contain intermediate zones. In 5 pegmatites this zone consists of muscovite—albite pegmatite surrounding one or more small discontinuous cores.

The Quartz Creek district has few well-zoned pegmatites. Those which are zoned commonly consist of only a wall zone and a core. The core units are usually irregularly distributed, discontinuous segments, and constitute only a small part of the pegmatite. Judging from the sizes and distribution of the core segments only a small proportion of pegmatitic liquid remained after consolidation of the wall zone. This crystallized in scattered areas as core segments.

Layered pegmatites.—Layered pegmatites make up approximately 7 percent of all pegmatites in the Quartz Creek district. Layering is most common in the thin dike-like lenticular and lenticular-branching types of pegmatites. Layering is not common in large irregular pegmatites, although a few of the thinner irregular bodies are layered.

Most of the layered pegmatites contain a perthite-rich hanging-wall unit and an albite-rich footwall unit (fig. 11). As an example, pegmatite No. 685 (pl. II) on the north side of Wood Gulch has a hanging-wall unit of albite (30 percent), quartz (20 percent), perthite (48 percent), and musco-vite (2 percent) and a footwall unit of albite (65 percent), quartz (15 percent), perthite (19 percent), and muscovite (1 percent). In a few pegmatites the hanging-wall unit has more albite than perthite but this unit always contains more perthite than the footwall unit. Pegmatite No.

1363 (fig. 11 and pl. II) is the only body that contains a higher proportion of perthite in what is believed to be the footwall unit; this pegmatite, however, is nearly vertical (81 degrees). The texture of the hanging-wall unit is coarser than that of the footwall unit, because perthite tends to form larger crystals than albite. Perthite forms in grains 0.5 to 3 inches in diameter and albite forms in grains 0.06 to 0.25 inch in diameter. In the albite-rich units the quartz grains are about the same size as the albite, but in the perthite-rich units they are nearly as large as the perthite.

Layered pegmatites with a perthite-rich hanging-wall layer and an albite-rich footwall layer have been described by Schaller _/ in the Pala

_/ Schaller, W. T., The genesis of lithium pegmatites: Am. Jour. Sci. 5th ser., vol. 10, pp. 271-274, 1925.

district, Calif.

The concentration of perthite as hoods in the upper part of zoned pegmatites is common in many districts, for example, the Keyes No. 1 pegmatite, Orange, N. H.; the W. T. Foster No. 1 pegmatite, Shelby, N. C., the Palermo No. 1, Grafton County, N. H., the Strickland-Cramer pegmatite, Portland, Conn., the Beecher Lode, Dyke Lode, Etta, Dan Patch, Hugo, and the Bob Ingersoll Dikes Nos: 1 and 2 of the Black Hills, S. Dak. _/ The

[/] Cameron, E. N., Jahns, R. H., McNair, A. H., and Page, L. R., Internal structure of granitic pegmatites: Econ. Geology, Mon. 2, pp. 44-45, 48, 1950.

perthite-rich hanging-wall layers in pegmatites of the Quartz Creek district appear to be the extreme development of perthite-rich hoods in thin lenticular bodies. Two of the lepidolite-bearing pegmatites, pegmatite No. 306 (Opportunity No. 4 claim) and pegmatite No. 452 (the Brown Derby

No. 1) are layered. Pegmatite No. 306 consists of an upper albite-quartz-perthite unit and a lower cleavelandite-quartz-lepidelite unit. The Brown Derby No. 1 pegmatite contains at least eight different units. Not all of these are present throughout the pegmatite and some form lenticular pods. The Brown Derby is described by Hanley _/ as having a border zone, wall

__ Hanley, J. B., Heinrich, E. W., and Page, L. R., Pegmatite investigations in Colorado, Wyoming and Utah: U. S. Geol. Survey Prof. Paper 227, pp. 69-71, 1950.

zone, possible intermediate zone, and a compound core of three different units. The Brown Derby has more mappable units than any other pegmatite in the region. Many of these layers are found in only certain parts of the dike and merge along strike with other units. The central part of the unit has an albite-quartz wall zone on both hanging-wall and footwall sides, but to the north the wall zone on the hanging-wall side disappears and the pegmatite becomes a layered pegmatite. Other layered pegmatites probably are incompletely developed zoned pegmatites.

The layered pegmatites are most abundant in (1) along the ridge just south of Quartz Creek in the southwestern corner of the district, (2) in the vicinity of the Brown Derby mine, and (3) along the western side of Big Gulch. Layered pegmatites are sparsely scattered among other types of pegmatites in the first two areas but are the dominant type in the third area. The layered type is almost absent in other areas. The distribution of the layered pegmatites suggests that their development is controlled by a particular set of conditions. The country rock in these areas is hornblende gneiss as in many areas that are void of layered pegmatites. The conditions under which these bodies cooled and crystallized probably controlled their formation. The original composition of the pegmatite liquid

may have been important, but probably was not the controlling factor, because there are different mineral assemblages in layered pegmatites, and because many unlayered pegmatites are identical to layered ones in mineralogy. The possibility of the layers being formed by replacement of preexisting pegmatite rather than by difference in crystallization history has been considered, but most of the perthite is in well-formed crystals, surrounded by later albite and quartz. The crystals are only slightly embayed and are not cut by veinlets of other minerals. The facts that the perthite layer, with only one possible exception, is on top, that the contact between the two layers is gradational, and that the layers may telescope gradually into a homogeneous unit do not seem to fit the picture of irregular replacement.

Pegmatites showing variation in composition along strike.—About 1 percent of the pegmatites have more than one unit, where the mineral composition of the unit changes along the length rather than across the pegmatite (fig. 12). In some lenticular-branching pegmatites, each branch has a different mineral assemblage. In lenticular pegmatites one end may be of one mineral composition and the opposite end the other, or the center of the pegmatite may be of one mineral composition and the ends a different mineral composition. In a few pegmatites one part may contain a core in addition to the layers across the pegmatite.

The dominant variation is from a unit rich in perthite to one in which perthite is less abundant or even absent; however, parts of some of the branching cleavelandite- and lepidolite-bearing pegmatites on the Opportunity No. 1 claim (No. 209, 213, 214, 215, and 216) are of this type.

Pegmatites showing variation in composition along strike are in lenticular, and lenticular-branching pegmatites and, in part, may represent multiple injections of pegmatite liquid.

Multiple pegmatites .-- Multiple pegmatites are formed by multiple intrusions so that the walls of the pegmatite formed by the second injection are tangent to that of the first. Thus, the various units have strikes which trend within a few degrees of one another. Branching and irregular pegmatite bodies may be the result of multiple intrusion of two or more pegmatite liquids into the same spot. In general, adjacent pegmatites probably are derived from liquid fractions of the same granitic magma and may be expected to consist of the same minerals in approximately the same proportions. It may be difficult, therefore, to distinguish between a branching and a multiple pegmatite. There are in the Quartz Creek district two pegmatites that have been formed by two separate injections of pegmatitic liquids. Pegmatite No. 251 is a lenticular-branching body with a wall zone and a thin core in each branch (fig. 13). The two branches join near the north end, and instead of the cores joining as cores do in a normal branching pegmatite, there are two parallel cores at the junction showing that it is a multiple pegmatite formed at slightly different times by two different injections. Pegmatite No. 216 on the Opportunity No. 1 claim is a north-trending albite-quartz-perthite pegmatite, which is cut by a northeasterly-trending body of perthite-quartz and cleavelanditequartz pegmatite (fig. 13). The north-trending part of the pegmatite mass is older than the northeasterly-trending branch. Though this pegmatite resembles a multiple pegmatite in that it is formed of two separate injections, the cross-cutting relationship proves it to be an earlier pegmatite cut by a later one.

Mineralogy

A total of 27 minerals has been found in the pegmatites of the Quartz Creek district. Perthite, plagioclase, and quartz are the essential min-

erals and form from 95 to more than 99 percent of most pegmatites. Only a very few pegmatites have units rich in muscovite, and this mineral cannot be considered an essential mineral of the pegmatites of this district. The common accessory minerals are considered to be those found in more than 10 percent of the pegmatites. These minerals, in order of their frequency, are: muscovite, garnet, biotite, magnetite, and beryl. The quantity of these minerals in any particular pegmatite is small; muscovite commonly ranges from 0.5 to 3 percent, garnet, 0.5 to 1 percent, biotite and magnetite, less than 1 percent, and beryl, a few small crystals.

The other 19 minerals are found in less than 3 percent of the pegmatites of the district and are considered as rare accessory minerals.

They also commonly amount to only a small fraction of a percent of the pegmatite. Table 3 lists all the accessory minerals giving the number of pegmatites which contain these minerals and the percent of the total pegmatites in which they are found.

Table 3.--Occurrence of accessory minerals in the pegmatites of Quartz Creek district

Mineral	Number of pegmatites in which mineral was observed	Percentage of pegmatites examined (1,803) in which mineral was observed
Muscovite	1,058	57.9
Garnet	965	52.9
Magnetite or martite	422	23.1
Biotite	357	19.6
Beryl	232	12.7
Tourmaline	48	2.6
Columbite-tantalite	29	1.6
Monazite	23	1.3
Lepidolite	17	0.9
Microlite	13	0.7
Chlorite	9	0.5
Topaz	8	0.4
Gahnite	8	0.4
Samarskite	7	0.4
Epidote	3	0.2
Apatite	7 3 3 2	0.2
Fluorite	2	0.1
Spodumene	1	0.06
Amblygonite	1	0.06
Allanite	1	0.06
Lithiophilite-trithylite		0.06
Betafite	1 1	0.06
Chrysocolla		0.06
Unknown	1	0.06

Plagioclase

Plagioclase occurs in all pegmatites of the Quartz Creek district and is the dominant mineral in most of them. It may form as much as 90 percent of all types of structural or mineralogic units. The plagioclase occurs in fine-grained sugary aggregates of equigranular grains, and in coarse platy crystals (cleavelandite). Cleavelandite has been found in 28 pegmatites in this district. Where used in this report, the term plagioclase refers only to the typical granular form, and the term cleavelandite refers to the platy form. Plagioclase commonly is abundant in (1) homogeneous pegmatites, (2) wall zones of zoned pegmatites, and (3) footwall layers of layered pegmatites. These units may contain more than 98 percent plagioclase and quartz. Cleavelandite, on the other hand, is restricted for the most part to central parts of the pegmatites.

The plagioclase ranges in size from less than 0.003 inch to about 1.5 inches across; the average size is about 0.12 inch. Crystal shape is usually not discernible, but the cleavage surfaces commonly are curved or warped. Twinning is visible only on the large pieces. Plagioclase is found in graphic intergrowth with quartz in a few places. Most of the plagioclase is white, but cream-colored, brownish, and pinkish shades are common. The plagioclase locally resembles the perthite in color, but can usually be distinguished by the warped

surfaces, twinning lamellae, lack of perthitic structure, and to a lesser extent by its occurrence in fine aggregates.

Cleavelandite is found in thin plates 0.003 to 0.006 inch thick and 0.5 to 4 inches in maximum dimension. The average plate is approximately 2 inches long, 1.5 inches wide, and 0.04 inch thick. The crystals are white and semi-transparent. The surface of the crystal is wavy and twin lamellae can be seen along the edges.

The lowest index (N₂₁) from the (OlO) cleavage flakes was determined on 439 specimens of plagioclase and on 17 specimens of cleavelandite from the Quartz Creek district. The index of the plagioclase ranged from 1.527 (Ab₉₉An₁) to 1.541 (Ab₇₄An₂₆) and of the cleavelandite from 1.528 (Ab₉₇An₃) to 1.530 (Ab₉₅An₅). The plagioclase has an average of 1.532 (Ab₉₃An₇) and cleavelandite 1.529 (Ab₉₆An₄). The plagioclase ranges from sodic albite to calcic oligoclase, and the cleavelandite is a sodic albite. Table 4 gives the results of these determinations, the type of country rock, and the type of pegmatite or pegmatite unit from which each specimen was taken.

Table 4.—The range of refractive index (Nol) of plagicalse and cleavelandite from all types of pegmatite units in the Quartz Creek district and its relation to different types of country rock.

Country rock							determ) of p			l refra	active				an	mber of d refrac	tive in	nations dex (N
Type of pegmatite	1																	
	527 1	528	1 520 1	570	571	572	1 533	1 534	1 535	1 536	1 537	1 538	1 530	1.540 1.	5117	1 528	1.529 1	530
or pegmatite unit . I.	201 1	020	10.)69	1.000	1)	10.))	10.000	+•.))+	10.000	1.9.90	1.001	1.7.70	1.7.77	1.)40 1.	.)-11	1.)20	10) 1	•))0
Hornblende gneiss																		
Homogenous pegmatite		16	41	39	11	11	8	6	8	4	10	5	2		1			
Zoned pegmatites		No. Vi		"														
Wall zone		2	25	11	9			2	2									
Intermediate zone(s)				1	í												1	
Core		1 2	5	5	1							1					6	2
Layered pegmatites					- A													
Lower layer			6	3	1				1									
Median layer			1	i														
Upper layer		1	7	1.5														
Pegmatites which change																		
in composition along																		
strike	1	2		1	1			100	1	1								1
Fracture filling				1				1										
Granite and hornblende gne	iss																	
or tonalite	-																	
Homogeneous pegmatites	2	3	10	9	6	1							t .					1
Zoned pegmatites																		
Wall zone			2	2 2	2	1		1										1
Core		1	1	2	3					100							1	
Layered pegmatites																		
Lower layer			2	1	2				1									
Median layer(s)			1															
Jpper layer			2	2	1		1		1				1					
Pegmatites which change																		
composition along stri	ike	1	1	3	1													
																1	2	

Table 4.—The range of refractive index (No1) of plagioclase and cleavelandite from all types of pegmatite units in the Quartz Creek district and its relation to different types of country rock.—Continued

Number of determinations and refractive Number of determinations index (No.) of plagioclase and refractive index (Ne) Country rock of cleavelandite Type of pegmatite 1.527 1.528 1.529 1.530 1.531 1.532 1.533 1.534 1.535 1.536 1.537 1.538 1.539 1.540 1.541 1.528 1.529 1.530 or pegmatite unit Fine- and coarse-grained granite Homogeneous pegmatites Zoned pegmatites Wall zone Core Quartz monzonite Homogeneous pegmatite 19 Zoned pegmatites Wall zones Intermediate zone(s) 1 Layered pegmatite Lower layer Upper layer Country rock unknown Homogeneous pegmatite 1 1

The refractive indices of plagioclase were determined for all internal units in each beryl-bearing pegmatite and for a selected number of pegmatites in each type of country rock. The refractive indices of the plagioclase in pegmatites with the hornblende gneiss and tonalite wall rocks averaged 1.530, and those in the quartz monzonite averaged 1.538. Refractive indices of plagioclase in pegmatites from the hornblende gneiss near quartz monzonite are in the same range as those in the quartz monzonite. The refractive indices of plagioclase in pegmatites decrease from the quartz monzonite area southward; the difference appears to be controlled by the regional distribution rather than composition of the country rock.

In the zoned pegmatites of the Black Hills _/ and other districts,_/

a systematic variation in the plagioclase has been noted from zone to zone with the anorthite content decreasing toward the core. Most pegmatites in the Quartz Creek district do not have the well-developed zones but when zoned show a rather large wall zone with scattered core segments. Table 5 shows that there is not much change in the indices of the 17 pegmatites that are composed of just wall zone and core; eight showed no change of index, five decreased slightly in index (0.001 to 0.002) from wall zone to core, while four increased in index (0.001 to 0.006). In two well-segregated pegmatites with intermediate zones there is a decrease in refractive index of 0.005 towards the core, indicating a decrease in anorthite content, which is in accord with previous work. In a well-segregated pegmatite the plagioclase probably will tend to show a systematic change of anorthite content toward the center; but when the zoning is poor, the plagioclase

Page, L. R., et al., Pegmatite investigations, 1942-1945, Black Hills, South Dakota: U. S. Geol. Survey Prof. Paper (in preparation).
 Cameron, E. N., et al., op. cit., p. 99, 1949.

will either have the same composition or have erratically distributed values.

Table 5.--Lower refractive index (N_{α_1}) of plagioclase in zoned pegmatites.

Pegmatite No. (pl. II)	Nal of plagioclase in wall zone	Naj of plagioclase in intermediate zone	Nal of plagioclase in core
174	1.529		1,528
279	1.531		1.529
289	1.529		1.530
451	1.529		1.529
453	1.529		1.529
454	1.529		1.529
455	1.529		1.529
456	1.529		1.529
535	1.535	1.529	
536	1.528		1.528
674	1.531		1.530
847	1.533		1,539
989	1.530		1.529
1002	1.531		1.530
1028	1.529		1.531
1044	1.529		1.530
1202	1.530		1.530
1402	1.537	1.532	
1666	1.529		1.529

A comparison was also made between the refractive indices of plagioclase from the hanging-wall layer and the footwall layer of two layered
pegmatites (table 6). As the hanging-wall layer is relatively rich in
perthite and the footwall layer relatively poor, it was thought that this
change in the alkali content might be reflected in the ratio of sodium to
calcium in the plagioclase. Of the eleven pegmatites investigated, six
showed no variation in index and five showed an increase of 0.001 to 0.002
in index from the hanging-wall layer to the footwall layer. As the limit
of accuracy of the index determinations is approximately 0.001, the results show a negligible change. The concentration of perthite and thus
the potassium content in the upper layer seems to have little effect on
the ratio of sodium to calcium in the plagioclase.

Table 6Lower	refractive	index	(N.)	of	plagioclase	in	layered.	pegmatites
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Pegmatite No. (pl. II)	Nal of plagioclase in hanging-wall layer	Nal of plagioclase in footwall layer
270 432 433 435 462 548 778 1004 1043 1105 1172	1.528 1.529 1.530 1.529 1.529 1.529 1.535 1.529 1.530 1.530 1.529	1.530 1.530 1.530 1.529 1.529 1.529 1.535 1.529 1.531 1.531

A comparison of the refractive indices of plagioclase in beryl-bearing and non-beryl-bearing pegmatites (table 11) shows that both have a wide range and that the non-beryl-bearing pegmatites have more calcic plagioclase.

Cleavelandite occurs in crystals commonly many times larger than the plagioclase, has a flat platy crystal habit compared to more equant grains of plagioclase; and is invariably white or bluish white whereas plagioclase is white, yellow, green or pink. These two varieties probably did not form under identical chemical and physical conditions. Cleavelandite occurs in 20 pegmatites in the Quartz Creek district and has been noted in homogeneous, zoned, and layered pegmatites, and pegmatites which vary in composition along the strike. In zoned pegmatites it is found in the intermediate zone in three places and in the core in sixteen. The tendency for cleavelandite to form in the central part of pegmatites has been noted in other parts of Colorado \int and in other districts in the United States \int .

Cleavelandite has distinctive mineral associations. In the Quartz

Creek district it is associated with quartz, lepidolite, microlite, beryl,

topaz, columbite-tantalite, perthite, muscovite, garnet, and tourmaline.

Many of these minerals are normally found with cleavelandite though the reverse is not always true. Cleavelandite is associated with lepidolite in 14 of the 17 lepidolite-bearing pegmatites, with topaz in 8 of the 8 topazbearing pegmatites, and with microditta in 12 of the 13 microlite-bearing pegmatites. The association of cleavelandite with lithium minerals and some of the rare accessory minerals has been noted in other districts. In the Tin Mountain pegmatite, Custer County, S. Dak., cleavelandite occurs in the core associated with spodumene, lithium mica, beryl, amblygonite, cassiterite, columbite-tantalite, apatite, microlite, and pollucite. In the Harding mine near Dixon, N. Mex. _/ cleavelandite in fracture fillings

_/ Adams, J. W., Personal communication.

is associated with purple muscovite, microlite, and spodumene. In the Rutherford and Morefield pegmatites near Amelia, Va. _/, cleavelandite is

_/ Glass, J. J., op. cit., pp. 761-763, 1935.

associated with cassiterite, manganotantalite, microlite, and zircon. The regional distribution of cleavelandite-bearing pegmatites (fig. 20) shows that with the exception of two bodies, all the deposits containing cleavelandite also contain lepidolite or are adjacent to pegmatites containing lepidolite. This suggests that some of the elements common to lepidolite and its associated minerals may be responsible for the formation of platy plagioclase. The elements that might promote the growth of cleavelandite are lithium (in lepidolite, zinnwaldite, amblygonite, and spodumene), rubidium or cesium (in lepidolite, muscovite, tourmaline, beryl, and pollucite), and fluorine (in lepidolite, fluorite, and topaz). Fluorine is more difficult to evaluate because it may be present as an essential

constituent, as in fluorite, topaz, or lepidolite, or it may occur undetected as a minor constituent in other minerals such as muscovite by substituting for the OH radical. Spectrographic analyses of both cleave-landite and albite from a number of pegmatites should be made to find whether small quantities of lithium, fluorine, or other elements are present in one type of plagioclase and not in another.

Perthite

All the potassium feldspar examined in pegmatites of the Quartz Creek district was white, cream-colored, or pink perthite; no orthoclase or microcline free of vein-like laminations of albite was noted. The albite laminae are thin, roughly parallel, and white. The albite is well twinned and the twinning is parallel in all laminae. The twin planes parallel the long dimension of the albite laminae and have the same orientation in a single perthite crystal.

Perthite occurs in most pegmatites and in some it is the predominant mineral. It is absent from some sodic-rich units, but forms as much as 93 percent of perthite cores. Sodic-rich pegmatites commonly contain less than 15 percent perthite; other pegmatites are in a large part graphic granite. Generally the perthite-rich pegmatites are most abundant in the northwest part of the district.

All pegmatite units contain perthite, but it is most abundant as graphic granite in homogeneous pegmatites, wall zones of zoned pegmatites, or as blocky perthite in cores of quartz-perthite pegmatite. The largest perthite-bearing pegmatites in the district contain graphic granite whereas graphic granite is not common in the small cores of zoned pegmatites. The hanging-wall unit of layered pegmatites is commonly rich in perthite.

The perthite is in crystals a quarter of an inch to 8 feet in maximum dimension. The crystals are largest in cores where they average 1.5 feet. Perthite crystals in the wall zone or in layered and homogeneous pegmatites are 2 to 3 inches in size. In fine-grained plagioclase-rich pegmatites the perthite crystals are smaller than in pegmatites where perthite is the dominant mineral. Graphic granite crystals are from a half to 4 feet in length and average about 80 percent perthite and 20 percent quartz. About a ton of graphic granite was crushed, quartered, and analyzed _/. This analysis (table 7) indicates that the soda is almost

__ Analysis obtained through the courtesy of C. A. Wemlinger, Vice President in Charge of Operations, Beryllium Mining Co., Inc.

entirely in the albite laminae of the perthite. Normative minerals _/,

/ Washington, H. S., U. S. Geol. Survey Prof. Paper 99, pp. 1162-1165, 1917.

calculated from this analysis, verify that there is little present other than quartz and feldspar and that the microcline molecule is 4.5 times as abundant as that of plagioclase.

Perthite forms blocky equidimensional crystals that are surrounded and veined by an aggregate of quartz, albite, and muscovite. In most places perthite is the first essential mineral to crystallize, but rarely it appears to be later than some or all of the associated minerals,

Table 7.--Chemical analysis 1/of graphic granite from the Bucky mine,
Quartz Creek district, Colorado

Oxide	Percent
SiO ₂	71.56
Al203	14,82
K ₂ 0	10.97
Al ₂ 03 K ₂ 0 Na ₂ 0	1.69
CaO	0,08
	0.01
Fe ₂ 0 ₃ Mg0 Cr ₂ 0 ₃	Trace
Cr203	None
2-7	99.13

1/ C. A. Parker, analyst.

Quartz

Quartz comprises 15 to 30 percent of all pegmatites in the district; the average is about 20 percent in homogeneous pegmatites, non-lepidolite-bearing layered pegmatites, and wall zones of zoned pegmatites. Although the ratio of perthite to plagioclase varies widely in these types of rock, the quartz content is nearly everywhere 15 to 20 percent. Many cores and fracture fillings are made up solely of milky quartz, whereas other cores and fracture fillings are mixtures of blocky perthite and quartz. The quartz in fracture fillings, intermediate zones, and cores is 10 to 100 percent of the unit.

The quartz is generally white to gray, although smoky varieties are found in a few pegmatites, usually as small oval blobs of from 1 to 10 feet. The smoky varieties are usually associated with radioactive minerals, for example, with microlite in pegmatites No. 215, No. 216, and No. 452, and with allanite in pegmatite No. 847. Small patches of smoky quartz have been found without visible radioactive minerals.

Quartz in most places fills interstices and forms veins in perthite crystals, indicating that it crystallized after the perthite. Rarely,

however, the reverse is true. The quartz associated with blocky perthite is in crystals 2 to 18 inches in size and is commonly slightly finer grained than perthite. In graphic granite, the quartz forms crude cuniform-shaped rods 0.03 to 0.25 inch thick and as much as 1.5 feet long.

Albite is interstitial to quartz and in places appears to vein it.

This relationship indicates that albite crystallized last, but in many places the mutual intergrowths suggest a contemporaneous age. Where quartz occurs solely with albite it forms crystals 0.03 to 0.5 inch in diameter; as the proportion of perthite increases in the unit the size generally increases. Muscovite and quartz in many places are intergrown and appear to have crystallized at about the same time.

Muscovite

Muscovite is found in about 60 percent of the pegmatites in the Quartz Creek district. On the east side of Quartz Creek it occurs in 85 percent of the pegmatites. On the west side the iron content of the pegmatites is higher, considerable magnetite is present, and biotite occurs in place of part of the muscovite. Muscovite is found in all types of internal units in the pegmatites and forms 0.5 to 3 percent of the rock; rare small pegmatites contain as much as 10 percent.

The muscovite is clear to green and individual sheets show black mineral staining. The larger pieces have reeves and "A" structure. Most of the muscovite, however, is about 0.25 inch in diameter, and commonly is intergrown with quartz. In only two pegmatites, the Buckhorn (No. 659) and the Bucky (No. 1574), are muscovite books more than 3 inches in size; books 1 foot in size occur in the Bucky pegmatite.

The muscovite is in both flat and curved forms. The flat variety is common in most rocks; the curved variety is found in 23 pegmatites, all on

the northwestern slope of Wood Gulch, where it occurs in a series of concentric shells 0.12 to 0.5 inch thick.

/ Winchell, A. N., Elements of optical mineralogy: pt II, p. 268, 1947.

Volk, G. W., Optical and chemical studies of muscovite: Am. Mineralogist, vol. 24, pp. 255-266, 1939.

composition of any sample of muscovite can be expressed in terms of these three end members. The end members are potassium muscovite (H₄K₂Al₆Si₆O₂A), phengite (H₆K₂(Fe,Mg)₂Al₄Si₆O₂A), and ferric iron muscovite (H₄K₂Fe₂Al₄Si₆O₂A). The refractive indices of muscovite increase with the proportion of the ferric iron muscovite in the mineral. The total amount of iron can not be ascertained by optical methods alone, however, as the iron may also be in the ferrous form in the phengite member. Specimens containing phengite and the potassium muscovite member have the same indices for equal amounts of the ferric iron muscovite. Information obtainable from refractive index determinations on the chemical composition of muscovite is therefore less useful than similar data on plagicalse and beryl. Volk \(\sqrt{} \) made

[/] Volk, G. W., op. cit., pp. 257-259, 1939.

²² chemical analyses and obtained the optical data on muscovite from various pegmatites scattered throughout the world. These analyses are in an area on the diagram midway between potassium muscovite, phengite, and 0 to 38 percent ferric iron muscovite. The median refractive index (N_B) was determined on 95 specimens of muscovite from the Quartz Creek district and ranged from $N_B = 1.585$ to 1.606 (table 8), indicating from 0 to 28 percent of the ferric iron muscovite molecule. It was thought originally that a

variation in the refractive indices, and thus in ferric iron content, might be found between units or layers. Table 9 shows the median refractive indices (N_B) of muscovite from the wall zone and core of seven zoned pegmatites. There is a small variation in the refractive indices but the variation is not constant, either in direction or amount. The median refractive indices of muscovite in layered pegmatites (table 10) show a small but unsystematic variation between the hanging-wall to the footwall layers. A comparison of median refractive indices of muscovite from pegmatites in various types of country rock was made. The lack of sufficient samples from pegmatites in the granite and quartz monzonite made this work inconclusive, but the variations are in the same range as those from pegmatites in the horn-blende gneiss or tonalite. Comparison was also made between the muscovite in various beryl-bearing and non-beryl-bearing units. The refractive indices of the muscovite in the beryl-bearing units were in the same range as those in the non-beryl-bearing units.

A ALL

Table 8.—Number and distribution of median refractive indices (Np) found in flat and curved muscovite

Median index	Number o	f specimens
(NB)	Flat muscovite	Curved muscovite
1.585	1	1
1.586	1	0
1.587	1	0
1.588	0	0
1.589	0	2
1.590	0	1
1.591	0	0
1.592	4	2
1.593	1	0
1.594	7	0
1.595	4	2
1.596	. 4	0
1.597	12	1
1.598	7	0
1.599	4	3
1.600	7	0
1.601	8	1
1.602	5	1
1.603	6	1
1.604	1	0
1.605	5	1
1.606	1	0

Table 9. - Median refractive index (No) of muscovite from zoned pegmatites

Pegmatite No.	Ng of muscovite in wall zone	Ng of muscovite in core
174	Anni pang pang pang	1.592
208	1.599	1.578 (zinnwaldite)
213	1.592	1.599
245	1.596	1.594
266	1.597	1.595
288	1.595	1.597
321	1.594	Sales applicated series \$440

Table 10. -- Median refractive index (No) of muscovite in layered pegmatites

Pegmatite No.	Ng of muscovite in hanging wall	Ng of muscovite in footwall
913	1.597	
927	1.604	1.600
937	1,597	
944	1.601	1.601
953	1.599	1.602
954	1.605	1.603
958	1,606	1.598
959	1.605	1.605
963	1.605	1.603
969	1.605	1,602
975	1.598	1.601
997	1.603	1.598
1132	1.601	1.601
1172	1.595	1.594

It was thought that the curved muscovite in the Quartz Creek district might also be a lithium mica, but median indices (Ng) determined for 16 specimens of curved muscovite show the same range of index and approximately the same distribution as the flat muscovite (table 8). Furthermore, the angle 2V of the curved mica (40 degrees) is much too high for a lithium mica. Several lithium micas were found, however, in making refractive index determinations on muscovite. These micas are colorless, flat, and associated with cleavelandite. Their median index ranges from 1.560 to 1.578 which is below that of the muscovite series. These specimens are in the zimmweldite range of the lepidolite series. There is no sure way to distinguish white lithium mica from muscovite in hand specimen. The lithium micas are more brittle, and the presence of cleavelandite should lead one to consider the possibility of lithium micas being present. A simple test to distinguish the two involves the use of a blowpipe: the lithium micas can be fused but muscovite can not.

Garnet

Approximately 55 percent of the pegmatites of the Quartz Creek district contain minor quantities of garnet. It is commonly in crystals less than 0.03 inch in diameter and may be overlooked easily. Garnet ranges in size from less than 0.01 inch to 1 inch in diameter, but crystals over 0.15 inch are rare. This mineral occurs in all the pegmatite units, but has a decided preference for the fine-grained plagioclase-rich parts, such as found in the footwall units of layered pegmatites, the wall zones of zoned pegmaties, and homogeneous pegmatites. It is found in crystals 0.20 inch and larger in the coarser-grained cores, but in most cores it is absent. Though garnet is widely distributed throughout the district, is constitutes only a trace to less than I percent of most pegmatites; in a few of the smaller ones it makes up as much as 1 percent of the rock. Garnet is erratic in distribution, and some parts of a pegmatite may contain several percent while others contain none. In the Bucky pegmatite (No. 1574) rock exposed in two pits contains several percent garnet, whereas in the same unit in other pits the mineral is absent. Brown garnet is conspicuous in "line rock", forming long thin bands which contrast with the white plagioclase-rich bands.

The garnet occurs singly or in clusters as light-brown, reddish-brown, and black euhedral crystals. Some crystals are black on the outside due to manganese staining, but others are black throughout the crystal. The garnet in many pegmatites is clear reddish brown with no manganese staining. One of the larger crystals is an intergrowth of garnet and quartz.

The garnet group may be divided into six members: almandite (Fe3Al2Si3O12), spessartite (Mn3Al2Si3O12), pyrope (Mg3Al2Si3O12), grossularite (Ca3Al2Si3O12), andradite (Ca3Fe2Si3O12), and uvarovite (Ca3Cr2Si3O12).

Ford _/, Fleischer _/, and Wright _/ have shown that garnet specimens do

/ Wright, W. I., The composition and occurrence of garnets: Am. Mineralogist, vol. 23, pp. 436-449, 1938.

not correspond usually to any single chemical type, but contain two or more molecules in solid solution. It was shown first by Ford _/ that the index

of refraction and specific gravity of a garnet depend in a simple and direct way on the chemical composition. He calculated the index of refraction and the specific gravity of 23 garnets from their chemical composition. These values agreed within less than 2 percent with those determined by direct measurement. The reverse process is not so simple, as a single determination of refractive index or specific gravity may correspond to several combinations of end members. It would be possible from a series composed of three different molecules to have a number of different combinations with the same index of refraction. The problem is somewhat simplified because all garnets are in one of two groups: the aluminum-bearing garnets (almandite, spessartite, and pyrope) and the calcium-bearing garnets (grossularite, andradite, and uvarovite). These two groups, as shown diagrammatically by Winchell _/, are miscible with each other only in limited amounts.

___ Ford, W. E., A study of relationships existing between the chemical, optical and other physical properties of the members of the garnet group:
Am. Jour. Sci., 4th ser., pp. 33-49, 1915,

[/] Winchell, A. N., Elements of optical mineralogy: pt. 2, p. 175, 1947.

Wright _ compiled 35 analyses of garnets from pegmatites and 18 from

granites which he converted into weight percent of the five common members of the garnet group, namely: almandite, spessartite, pyrope, grossularite, and andradite. His conclusions from studies of garnet from pegmatites and other types of rock are: (1) that there is a remarkable constancy of one variety of garnet in each rock type, and (2) that spessartite and almandite constitute 85 to 90 percent of the molecules from pegmatites and granites. Thus, if one of the major constituents is known, the other can be estimated within a limit of error of 5 to 15 percent. Winchell \int has compiled data

✓ Winchell, A. N., op. cit., pp. 179-181, 1947.

by Ford and others into several diagrams from which, if the garnet group is known, and the specific gravity and index of refraction have been determined, a general composition in terms of the garnet molecules can be derived.

Indices of refraction were determined on garnet from 15 widely scattered pegratites in the Quartz Creek district. Specific gravity was not determined, but all specimens were qualitatively tested and found to contain manganese. All the indices of refraction are between 1.810 and 1.820 with many specimens having refractive indices about 1.815. The indices of refraction show small variations, but in general the garnet of this region is remarkably similar in index and composition. The proportion of almandite and spessartite can be roughly evaluated by neglecting the small percent pointed out by Wright to be taken up by the other garnet molecules and assuming that the mineral to be made up only of spessartite and almandite. In this case the garnet would range from 67 percent spessartite, 33 percent

almandite (N = 1.1810) to 33 percent spessartite, 67 percent almandite (N = 1.820). Most of the values would be closer, however, to 50 percent spessartite and 50 percent almandite (N = 1.815).

Garnet is associated with all the common and almost all the rare pegmatite minerals. It does have, however, a tendency to occur more abundantly with fine-grained plagioclase. In pegmatite units that are perthite-rich,
garnet if present commonly will be associated with the plagioclase.

Magnetite and martite

Magnetite, commonly altered to martite, is widespread in minor quantities and is found in approximately 20 percent of the pegmatites. Most pegmatites contain only a few scattered crystals, but several of the smaller pegmatites have about 1 percent.

Magnetite and martite are dull to steely black in color and rarely form well-developed octahedra. Almost all specimens, however, have good octahedral (111) parting which easily distinguishes this mineral from columbite-tantalite. The mineral ranges in size from grains less than 0.10 inch in size to round masses as much as 3 inches in diameter.

Magnetite and martite are found as an accessory mineral in the feldspathic pegmatites, but are not found in any of the lepidolite-bearing
units. They are in both the perthite-rich and the albite-rich pegmatites,
and are one of the few accessory minerals found in graphic granite. The
distribution of magnetite and martite is usually erratic; a few small areas
in the pegmatite may contain 1 or 2 percent and the rest of the pegmatite
only a trace. They are associated commonly with perthite, albite, quartz,
and biotite, and in a few places with garnet. Few beryl-bearing pegmatites contain either magnetite or martite; the two minerals are nowhere
adjacent to each other. Muscovite is nowhere associated closely with the

magnetite although they may both be in the same pegmatite, whereas biotite is closely associated with magnetite or martite.

The association of magnetite with biotite but not with muscovite is easily explained. Those parts of the pegmatite with sufficient iron to form magnetite also had sufficient iron to form biotite; those parts free of iron would contain muscovite in place of biotite.

Biotite

Biotite is found in almost 20 percent of the pegmatites on the west side of Quartz Creek, but is found in only 6 percent of the pegmatites on the east side. In most of the pegmatites biotite forms considerably less than 1 percent of the rock; in a few of the smaller pegmatites it forms several percent.

Biotite is dark to greenish black and occurs in widely scattered blades from a fraction of an inch to 8 inches in maximum dimension; in most pegmatites the blades are 0.25 to 0.5 inch. The larger blades usually occur in small areas and may be either restricted to core segments or small patches in the otherwise uniform homogeneous pegmatite or the wall zone of a zoned pegmatite.

The median refractive index (No) of seven specimens ranges from 1.636 to 1.671. Not only do the refractive indices vary from specimen to specimen but also in different parts of the same book. Much of the biotite is partly altered to chlorite and the variation in refractive index depends on the extent to which the biotite has been altered. These median refractive indices indicate that the biotite approximates siderophyllite \int in

composition and is high-iron rather than high-magnesium biotite.

Biotite was found in part of the lepidolite-bearing pegmatites, but it is not in the same units as lepidolite. It is common in both perthite- and albite-rich pegmatite and is one of the few accessory minerals in graphic granite. Biotite is commonly associated with magnetite or martite; musco-vite is found only in the magnetite-free part of these pegmatites.

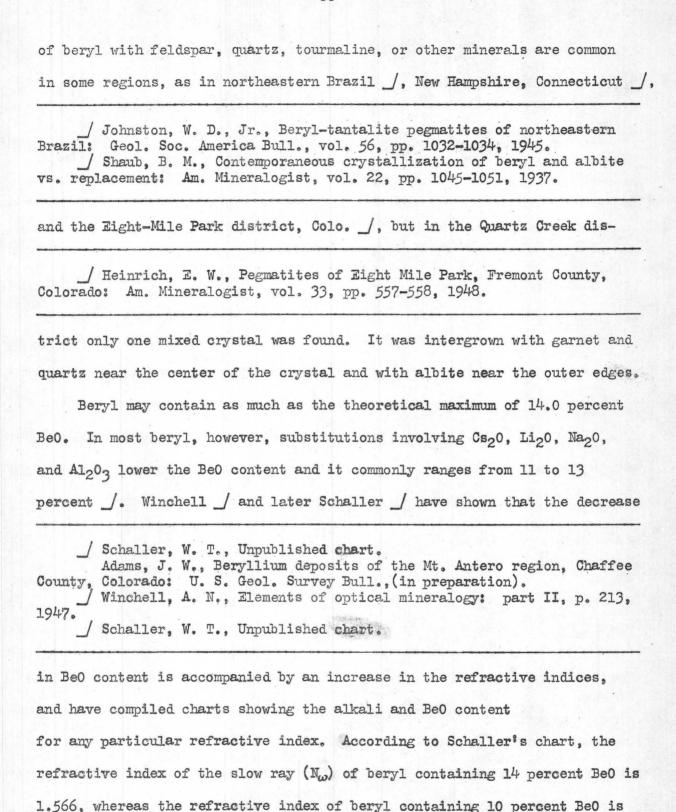
Beryl

Beryl is found in 232 pegmatites and is widely distributed; in most pegmatites in the Quartz Creek district there are only a few small crystals.

Beryl is found in all types of pegmatites and pegmatite units: homogeneous pegmatites; core, intermediate, and wall zones of zoned pegmatites; various layers of layered pegmatites; and units of pegmatites that differ in composition along strike.

Beryl may be brown, white, gray, greenish white, pale green, greenish gray, or pale blue green. The white, brown, and greenish white beryl is the most common, and it is difficult to distinguish from feldspar in many exposures. The beryl crystals range from 0.006 inch to 2 feet in diameter. In fine-grained albite-rich pegmatites the beryl crystals are 0.10 to 0.25 inch in diameter, but larger crystals occur in the coarser-grained intermediate zones and cores. Although beryl was found in a higher percentage of the albite-rich units than in perthite-rich ones only the latter contained beryl in pieces large enough to be hand cobbed. The average beryl crystal is approximately twice the size of the albite and about half the size of the associated perthite grains.

Beryl occurs as subhedral to euhedral hexagonal crystals; tapered crystals are rare except at the Bucky pegmatite (No. 1574). Intergrowths



The refractive index of the slow ray (N_{ω}) , determined for 183 beryl specimens from various units, ranges from 1.573 to 1.585 and averages 1.578.

1.600.

These determinations are compiled in table 11 together with the determinations of the minimum refractive index on cleavage plates of the associated plagioclase. The table is divided according to country-rock types and subdivided according to the type of pegmatite and internal structure. This table shows that the refractive index of beryl, and therefore the composition, varies irregularly in the different types of pegmatites and internal units. There appears to be no correlation between type of country rock and the refractive index of beryl. Only two specimens of beryl were obtained from pegmatites in the quartz monzonite and six from pegmatites in fine— and coarse—grained granite.

It has been noted in the Black Hills _/ and other districts _/ that

there is a systematic increase in the alkali content of beryl from the wall zone inward toward the core. Similar data on beryl from zoned pegmatites from the Quartz Creek district is rather meager because a large part of the beryl found was from a single zone of a zoned pegmatite. The refractive indices of beryl from different units of zoned pegmatites are compiled in table 12 together with the minimum refractive index of albite for comparison. Most pegmatites have only small cores and are zoned poorly. A difference of 0.001 is all that is noted from wall zone to core of the more simply zoned pegmatites. The refractive indices of the slow ray of beryl from three pegmatites that contain beryl in an intermediate zone as well as either a core or a wall zone differ as much as 0.003 between the intermediate zone and either the core or wall zone. This increase inward in alkali content is in accord with the findings of previous workers.

Table 11.-- The number of refractive index determinations of plagioclase and beryl, their relation to different types of country rock, and types of pegmatites or units.

Country rock and type of unit			tive in f plagi		cleavage	Highes	t refra	ctive i	ndex on	cleava	ige plat	es of l	eryl
	1.528	1.529	1.530	1.531	Other	1.575	1.576	1.577	1.578	1.579	1.580	1.581	Other
Hornblende gneiss and tonalite													
Homogeneous pegmatite	13	16	21	6	1.532-5 1.533-3 1.534-1	14	7	10	17	13	2	3	1.573-1 1.582-2 1.583-1 1.585-1
Zoned pegmatites													,, -
Wall zone	1	18	6		1.534-1		3	3	1	1	1		
Intermediate zone(s)	1	1	1	1						1 4	1		
Core	1	7	3	1		3	2	4	7	4	2		1.582-2
Layered pegmatites													
Lower layer		5	3	1				1		1			
Median layer(s)		1							1 2				
Upper layer	1	6	5	1		1	2	3	2	1			
Pegmatites which change													
composition along strike	2	1	3	1	1.527-1					2			
Fracture filling			1		1.534-1		2						
Granite, and hornblende gneiss and tonalite													
Homogeneous pegmatite	3	8	9	6	1.527-2 1.532-1	2	2	4	6	3	4	2	1.574-1 1.584-1 1.585-1
Zoned pegmatites													1.00
Wall zone		2	2	1							1	. 1	
Intermediate zone(s)		1			1.533-1		1						
Core	1	2	2	3	1.539-1	1	2		3	3	2		
Layered pegmatite													
Lower layer		2	2	2	1.535-1		1	1					
Upper layer		2	1		1.535-1			1	1				
Pegmatites which change					222								
along strike	2	3	2	1				2	2		1		1.574-1

Table 11.-- The number of refractive index determinations of plagioclase and beryl, their relation to different types of country rock, and types of pegmatites or units.-- Continued

Country rock and type of unit		ctive in		cleavage	Highes	st refra	active i	index on	cleava	ige plat	es of b	eryl
				Other	1.575	1.576	1.577	1.578	1.579	1.580	1.581	Other
Fine- and coarse-grained granite Homogeneous pegmatite	3	1	1			1				1	1	1.583-2 1.584-1
Quartz monzonite Zoned pegmatites Wall zone Core				1.537-1 1.532-1 1.535-1				1		1		1.70-1
Country rock unknown Homogeneous pegmatites		1		1.532-1	1	1						

Table 12.—Refractive indices of plagioclase and beryl from beryl-bearing units of zoned pegmatites

k.,	Wall zo:	ne	Intermediate	zone	Core	
Pegmatite No.	Nai of plagioclase	Nw of beryl	N _{&l} of plagioclase	Nw of beryl	N _{≪l} of plagioclase	Nw of beryl
174 250	1.529	1.580			1.528	1.580
279 289 436	1.531 1.529 1.529	1.576 1.578			1.529	1.578
454 455 5 3 5	1.529 1.529 1.534	1.577	1.529	1.580	1.529	1.578
847 989 1002 1025 1044	1.533 1.530 1.531 1.529 1.529			1.576	1.539 1.529 1.530 1.531 1.530	1.579 1.580 1.578 1.578 1.578
1202 1402	1.530		(1) 1.532 (2) 1.535 (1) 1,528	1.578	1.530	1.578
1574			(1) 1.528 (2) 1.530	1.579	1.528	1.582

Refractive indices of beryl and plagioclase from layered pegmatites have not been published. iTable 13 gives the refractive indices of these two minerals in 12 two-layered pegmatites. These pegmatites contain an upper layer of perthite-rich rock and a lower layer of albite-rich rock. Most of the beryl is found in the coarser-grained upper layer; in two pegmatites it occurs in both layers. In these two pegmatites the beryl showed essentially no change in refractive index between units.

Beryl is associated with plagioclase, quartz, perthite, muscovite, garnet, lepidolite, tourmaline, topaz, microlite, tantalite, monazite, gahnite, and biotite. Beryl is not associated with any one of these pegmatite minerals to the exclusion of the others. It has not, however, been found in graphic granite pegmatite. Beryl usually is in clusters or groups of crystals; in many pegmatites only three or four closely spaced crystals

are found. In pegmatite No. 279, 35 crystals of beryl were found in an area about 2 feet square. This was the only beryl noted, although this branching pegmatite exceeds 720 feet in length. Many other pegmatites have a similarly spotty distribution of beryl.

Table 13.—Refractive indices of plagioclase and beryl from beryl-bearing layered pegmatites

	Lower lay	er	Upper lay	er
Pegmatite No.	N _{al} of plagioclase	Nw of beryl	N _{wl} of plagioclase	Nw of beryl
270 417 432 433 435 462 548 778 985	1.530 1.529 1.530 1.530 1.529 1.529 1.529 1.535	1.577 1.578	1.528 1.530 1.529 1.530 1.529 1.529 1.529 1.535 1.535	1.577 1.576 1.577 1.575 1.578 1.576
1004 1105 1172	1.529 1.530 1.531	1,579	1.529 1.531 1.529	

Tourmaline

Tourmaline is relatively rare in the Quartz Creek district, and has been found in only 48 of the 1,803 pegmatites studied. In many pegmatite areas tourmaline is present in most pegmatites, but the Quartz Creek district is distinctive for its lack of tourmaline and the low boron content of its pegmatites. Except in the lithium-bearing pegmatites, only a few grains occur in each pegmatite. Some units in the lithium-bearing pegmatites contain as much as 3 percent tourmaline.

The tourmaline is black, dark green, blue, light green, and pink. The green, blue, and pink varieties are found only in the lepidolite-bearing pegmatites; the black variety occurs in both lepidolite- and non-lepidolite-

bearing pegmatites. Of the 48 tourmaline-bearing pegmatites 38 contain only the black variety. It occurs in subhedral to anhedral crystals; commonly the m(1010) and a(1120) prism faces are the only faces developed. In many places it is in small pods of coarse-grained quartz or quartz-perthite pegmatite in an otherwise homogeneous body. Black tourmaline has been found associated with quartz, perthite, albite, muscovite, beryl, garnet, cleavelandite, biotite, monazite, columbite-tantalite, and gahnite, but not with lepidolite and topaz.

The black tourmaline was found only in outer zones, completely free of lepidolite; in many zones it is restricted to the extreme hanging-wall or footwall part, Dark-green tourmaline was found either in zones containing lepidolite or the adjacent zones. It is common in the outer part of lepidolite-bearing units and the inner part of the adjacent unit. It is nowhere in contact with lepidolite, but occurs in the cleavelandite-quartz part of the zones. Pink and pale-green tourmaline occur adjacent to lepidolite in lepidolite-bearing units. These two varieties of tourmaline commonly occur together; the pink variety is more abundant. In the Brown Derby No. 1 pegmatite (No. 452) these varieties are together as "watermelon" tourmaline, in crystals which have a pink core and light-green rim. The dark blue variety of tourmaline is not present as individual crystals, but forms massive wavy bands in lepidolite-bearing pegmatites where it occurs in part with the black tourmaline and in part with the dark-green tourmaline. There is commonly a thin band of small garnets in the center of the blue bands. Figure 15, a sketch of part of pegmatite No. 453, shows the relation of various colored tourmalines to the pegmatite units.

The indices of refraction vary as well as the color. Table 14 gives the higher refractive index (N_{ω}) of 19 tourmaline specimens; eleven are of

black tourmaline and range from 1.652 to 1.664. The two specimens with the lowest refractive indices are from the outer edges of lithium-bearing pegmatites. Three dark-green tourmaline specimens have refractive indices of 1.646 and 1.647, all lower than black tourmaline. The black variety grades into the dark-green variety and it is to be expected that, if enough determinations of indices were made, a gradational series would be formed, with the green varieties having the lower refractive indices. Two pink tourmaline specimens from different pegmatites have refractive indices (No) of 1.643 and 1.637. A pale-green tourmaline and a pink tourmaline from adjacent areas in pegmatite No. 452 have a refractive index for the slow ray of 1.637; the pink core and the pale-green rim of a "watermelon" tourmaline had a refractive index for the slow ray of 1.634. pale-green and pink varieties appear to vary in composition and variation can not be correlated with color. The different colors may be caused by either the presence of a minor element that does not affect the refractive indices, oxidation or reduction of an element in different states of oxidation, or to a slight rearrangement of the molecular structure of tourmaline. The sequence from black tourmaline on the outer parts of lepidolite-bearing pegmatites to pale-green and pink tourmaline in the lepidolite-bearing part can be correlated with a progressive change in refractive indices, but the pale-green and pink crystals which grew together in the same environment without detectable changes in index can not.

Table 14. -- Refractive index (Nw) of tourmaline, Quartz Creek district

Pegmatite No.	Pegmatite Name	Color	И
205		Black	1.657
215	Opportunity No. 1	Black	1.652
231		Black	1.657
251		Black	1.657
306	Opportunity No. 4	Black	1.655
311		Black	1.663
1238		Black	1.657
1238		Black	1.659
1278		Black	1.6555
1322		Black	1.664
1607		Black	1.664
215	Opportunity No. 1	Dark Green	1.646
306	Opportunity No. 4	Dark Green	1.647
306	Opportunity No. 4	Dark Green	1.646
306	Opportunity No. 4	Pink	1.643
452	Brown Derby No. 1	Pink	1.637
452	Brown Derby No. 1	Pale Green	1.637
		(Pale Green	
452	Brown Derby No. 1	rim)	1.634
-72	220 WIT 2020 Y 100 T	(Pink center)	

Early work has shown that change in color in tourmaline commonly varied with the density of the crystals _/, the axial ratio a:c _/, and the re-

fractive indices. Color also was shown to vary with the chemical composition by more recent spectrographic work _/; Carobbi and Pieruccini _/ con-

clude from their studies that the pink color is caused by manganese with lithium and cesium. In the Quartz Creek district, manganese is present in most units but appears to be concentrated in the outer zones as indicated

J Warner, T. W., Spectrographic analysis of tourmalines with correlation of color and composition: Am. Mineralogist, vol. 20, pp. 531-536, 1935.

_/ Carobbi, G., and Pieruccini, R., Spectrographic analysis of tourmalines from the island of Elba with correlation of color and composition:

Am. Mineralogist, vol. 32, pp. 121-130, 1947.

by the higher concentration of garnets. The presence of manganese in a unit does not insure that the tourmaline will be pink, though the pink variety usually is high in manganese. Shainin _/ had spectrographic analyses made

of the minor base elements in four tourmaline specimens (table 15). These analyses show that the greatest amount of manganese (0.5 percent) occurs in a pink specimen and that all specimens contain manganese but that another pink specimen contains no more manganese than the green and blue specimens.

Table 15.—Spectrographic determination of minor elements in tourmaline from Maine 1/

Color	Mn	Ti	Ga	Sn	Pb	Zn
Deep blue	.2	.01	.02	.01	<.001	.5
Green	.2	.005	.02	.05	.02	.1
Pink	.5	<.001	.02	.01	-	.2
Deep pink	.2	<.001	.01	.05	.05	.02

1/ Analysis made in the Investigations Section of the Geochemistry and Petrology Branch of the U, S. Geological Survey for Vincent Shainin. Janet Fletcher, Analyst.

Lithium and cesium are both more abundant in the lepidolite units, lithium is a major component and cesium a minor component. Stevens _/ in an

_____ Stevens, R. E., New analyses of lepidolites and their interpretation: Am. Mineralogist, vol. 23, p. 615, 1938.

article on lepidolites gives the analyses of 17 lepidolites from widely scattered districts in which Cs₂0 ranged from 0 to 0.67 percent; the average was 0.23 percent. These analyses show that cesium is concentrated in appreciable amounts with the lithia, and that the two are comparatively abundant in those parts of the pegmatite where the pink and pale-green varieties of

tourmaline occur. The coloring elements are difficult to determine because of the large number that are present in minor amounts in tourmaline.

Tourmaline acts as a scavenger and takes into its structure small quantities of a great variety of elements.

found in tourmaline: B, Si, Al, Ti, Fe, Ni, Mg, Mm, Cu, Ca, Na, K, Ba, Li, Sr, Cs, Be, V, Ta, Sc, Sn, Ce, Ga, Pb, Zn, Cr, Co, Se, Ag, and Sb. Some of these elements are reported in tourmaline from only one area, while others are almost universal. The presence of some of these elements depends on whether or not they were available in the pegmatitic liquid at the time the tourmaline was formed.

Many elements appear to have little or no effect on the color of tourmaline as they are present in some colored tourmalines and absent in others.

Other elements may be responsible for a change in color only when found together with some other elements.

In resume, all tourmalines in the Quartz Creek district other than the black variety are in the lepidolite-bearing pegmatites. There is a gradual color change from black through blue and dark green to pink and light green paralleled by a change in refractive indices as the lepidolite-bearing parts of the pegmatite are approached. Changes in the concentration of a group

of alkalies, such as lithia, cesia, and probably others parallel the changes in color of the tournalines. The amount of alkalies probably is responsible for the color change and these elements probably have a greater affinity for the tournaline structure than many metallic elements. Iron and manganese are usually equally available in most pegmatite units. When iron is allowed into the structure it helps to darken the mineral to black or dark green, but the development of the lighter shades demands the presence of lithia and other alkalies. The various shades of green and blue are probably caused by the presence of various other elements.

Columbite-tantalite

Columbite-tantalite has been found in 29 pegmatites of the Quartz Creek district. It occurs in homogeneous pegmatites; wall zone, intermediate zone, and core of zoned pegmatites; layered pegmatites; and in parts of pegmatite which show variation along strike. These columbite-tantalite-bearing pegmatites are widely scattered over the entire district.

Columbite-tantalite is found in only a few crystals in most pegmatite units except in pegmatites No. 1234 and No. 452 (Brown Derby No. 1). In the latter pegmatite it makes up 1.4 percent of the rock in a small unit about 20 feet long and 1 foot wide.

In the Quartz Creek district the columbite-tantalite is black with a dull to lustrous surface. It has a black to brown streak. The tabular crystals range from the thickness of a sheet of paper to 1 inch and are from a fraction of an inch to 4 inches long. The crystals are usually subhedral to euhedral with the brachipinacoid, b(010), forming the tabular faces present in most specimens. Other faces which were noted on some of the columbite-tantalite crystals are: a(100), d(110), g(130), k(011), and u(111).

The columbite-tantalite, (Fe,Mn)(Cb,Ta)₂0₆, series is one of complete gradation between iron and manganese, and columbium and tantalum. Members of this series are divided on purely arbitrary standards with the columbite consisting of that part of the series where columbium exceeded tantalum in amount, and the tantalite part of the series where tantalum is in excess. A secondary division is made in these two main divisions by naming the mineral ferrocolumbite or ferrotantalite if the ratio of iron to manganese is greater than 3:1, and manganocolumbite or manganotantalite if the ratio of manganese to iron is in excess of 3:1. The specific gravity, the streak \(\sqrt{} \), and prob-

De Almeida, S. C., Johnston, W. D., Leonardos, O. H., and Scorza,

Hanley, J. B., et al., op. cit., p. 71, 1950

E. P., The beryl-tantalite-cassiterite pegmatites of Paraiba and Rio Grande do Norte, Northeastern Brazil: Econ. Geology, vol. 39, p. 218, 1944.

ably certain other physical properties vary with the columbium and tantalum content. Because chemical analysis of these two elements is expensive, the approximate composition is obtained by specific gravity determinations and reference to charts that related the specific gravity to the columbium—tantalum ratio. The ratio of iron to manganese has only a minor effect on the change of gravity, and the higher the specific gravity the higher the tantalum content. Table 16 gives the specific gravity of 8 specimens from the Quartz Creek district as determined on a Jolly balance. These specimens range from an almost pure manganocolumbite (specific gravity 5.0 and 5.1) to a columbium—rich tantalite (specific gravity of 6.7). As only the latter specimen falls in the tantalite field, this district appears to be one that contains columbite almost to the exclusion of tantalite. Hanley / gives the

specific gravity of a piece of columbite-tantalite from pegmatite No. 452 (Brown Derby No. 1 claim) as 5.61 and the chemical composition of 72 percent Cb₂O₅ and 6 percent Ta₂O₅.

The specimen on which the chemical work was done was collected by Eckel \mathcal{L} ,

_/ Eckel, E. B., op. cit., p. 244, 1933.

and was a different specimen from a different pegmatite unit from the Hanley specimen. Because the columbium-tantalum ratio commonly varies from zone to zone and evidently does on the two specimens used, the agreement is not good between the composition obtained from the specific gravity and that given by Eckel.

Table 16. -- Measurements of specific gravity on columbite-tantalite

Pegmatite Number	Internal Unit	Specific gravity
205	Core	6,1
205	Core	6.3
245	Core	
452	Layer	5.8
1234	Wall zone	5.7 5.8 5.0
1234	Wall zone	5.1 6.7
1557	Core	6.7
1574	Intermediate zone	6.0

Columbite—tantalite is found in direct contact with the following minerals in one or more pegmatites: quartz, albite, perthite, beryl, musco—vite, monazite, biotite, tourmaline, and gahnite. It also has been found in the same zones, but not in direct contact with: garnet, topaz, micro—lite, martite, and lepidolite. Though it is associated with almost all the pegmatite minerals, there are three associations which are most common in the Quartz Creek district: (1) with massive quartz, (2) with cleave—landite or cleavelandite and quartz, and (3) with feldspar (either perthite or plagioclase) and monazite. That its association with monazite is not pure happenstance can be seen in that 9 of the 24 monazite—bearing pegma—

tites, or 37 percent, contain columbite-tantalite (fig. 19).

Monazite

Monazite, (Ce,La,Nd,Pr)PO4, is found in 24, or approximately 1.5 percent of the pegmatites. It occurs in homogeneous pegmatites; cores, pods, and intermediate zones in zoned pegmatites; and in layered pegmatites. In three pegmatites, namely the Brown Derby No. 1 (No. 452), the Black Wonder (No. 847), and the Bucky (No. 1574), monazite is found in more than a half-dozen crystals. A unit 20 feet long and 1 foot wide at the Brown Derby No. 1, pegmatite No. 452, contains approximately 212 percent. Of two localities in the Black Wonder pegmatite (No. 847), one is worth special attention in that it contains 0.05 percent monazite in an intermediate zone of plagioclase-muscovite-quartz pegmatite surrounding a quartz pod. This intermediate zone is about 4 feet thick, and the quartz pod is approximately 15 feet long and 6 feet wide. In the Bucky pegmatite (No. 1574) the monazite occurs erratically in the mica zone around the quartz core.

Monazite occurs as euhedral, dark-red to clove-brown crystals, that range in size from 0.25 inch long, 0.12 inch wide, and 0.03 inch thick to 2 inches long, 1.5 inches wide, and 0.5 inch thick. Most of the larger specimens come from the Brown Derby No. 1. Crystal forms identified include the a(100), m(110), n(120), v(111), r(111), x(101), c(001), and h(305) faces. The crystals are usually flattened parallel to the a(100) face, and some of them are also twinned parallel to this face.

The specific gravity of the monazite varies from 5.0 to 5.6, as determined by the Jolly balance.

Optically, the monazite is colorless to yellow, with high bire-

fringence. The lower index of refraction, (N_d), ranges from 1.78 to 1.80, averaging about 1.79. Table 17 shows that the specific gravity and lower indices (N_d) do not vary with a consistent relationship.

In the Quartz Creek district monazite is associated with quartz, albite, perthite, muscovite, columbite-tantalite, gahnite, biotite, and garnet. It usually is found, however, in a feldspar-rich part of the pegmatite, and commonly produces a red stain in the feldspar immediately adjacent to it. Of the 24 pegmatites containing monazite, 9 also contain columbite-tantalite.

Table 17.—Lower index of refraction (N_{∞}) and the specific gravity of monazite

Pegmatite Number	Internal unit	Index of refraction (Na)	Specific gravity
452	Pod	1.80	5.3
290	Pod	1.79	5.1
847	Pod	1.78	5.1
847	Intermediate zone	1.78	5.6
997	Footwall layer	1.79	5.3

Lepidolite

Lepidolite is found in 17 pegmatites, comprising homogeneous pegmatites, core, and intermediate zones of zoned pegmatites, interior layers of layered pegmatites and in parts of pegmatites which show variation in composition along strike. Thus, no particular type of pegmatite seems to be favored. In the limited number of zoned pegmatites it appears to be commonest in the central parts.

The lepidolite is white, lilac, or various shades of purple; lilac to purple varieties are most common. It has three forms: (1) fine-grained aggregates with individual sheets less than 0.25 inch in diameter,

(2) large platy books 2 to 10 inches in diameter, and (3) curved concentric books, 0.5 to 2 inches across (pl. X). The large-plate lepidolite is found in 5 of the 17 pegmatites and form a group between the Brown Derby No. 1 dike (No. 452) and the Brown Derby No. 5 dike (No. 535), a maximum distance of 2,200 feet. Only three pegmatites on the Brown Derby No. 1 claim (No. 452, 454, and 457) have curved lepidolite. The plate and curved lepidolite are either purple or lilac. An analysis is reported by Stevens _/ on large

plates of pale-purple lepidelite from Ohio City. The sample probably came from either dikes No. 452, No. 453, or No. 454 (Brown Derby No. 1 claim), as these dikes were the only ones that had been developed at the time containing lepidolite in large plates. The analysis follows:

Li ₂ 0	5.05	Mg0	None
Si02	49.58	MnO	2.78
A1203	23.87	TiO2	0.06
K20	10.14	H_0=	0.51
NaO	0.57	H ₂ 0+	1.22
CaO	None	${\tt F}^{\sim}$	7.49
Rb20	1.62		103.19
Cs ₂ O	0.09	Less $0 = \mathbf{F}$	3.15
FeÕ	0.21 1/		100.04

1/ Total iron reported as FeO.

The formula for this lepidolite as determined from the analysis is $K_{\mu} \text{Li}_7 \text{Al}_2 \text{Si}_{15} \text{O}_{40}(\text{F,OH})_8$. Hanley $\sqrt{\text{reports that the physical and}}$

[/] Stevens, R. E., New analyses of lepidolites and their interpretation: Am. Mineralogist, vol. 23, p. 615, 1938.

_/ Harley, J. B., Heinrich, E. W., and Page, L. R., Pegmatite investigations in Colorado, Wyoming, and Utah: U. S. Geol. Survey Prof. Paper 227, p. 72, 1950.

optical properties of all the lepidolite from the Brown Derby claim (pegmatites No. 452, No. 453, and No. 454) are similar and the only chemical difference is that the manganese content of the large plates is slightly



Lepidolite from left to right: fine-grained aggregate, curved plates, and large plates.

higher. Winchell / in a triangular diagram shows that the median index of

____ Winchell, A. N., Elements of Optical Mineralogy: Part II, 3d ed., p. 271, 1927.

the lepidolite group of mica increased from the polythionite (H2K2Li3Al3Si8O24) and lepidolite (H4K2Li3Al5Si6O24) end members to the protolithionite end member, (HukaliFeuAlaSi6024). This increase of refractive index is in general paralleled by an increase of iron and a decrease of lithium and can be used to determine the approximate Ligo content. The median index (Ng) of 14 lepidolite specimens from six different pegmatites (table 18) ranged from 1.555 to 1.578. Six specimens are white, or white with just a tinge of lilac, and these specimens have the highest indices (1.560, 1.564, 1.564, 1.565, 1.575, and 1.578). The two specimens with the 1.575 and the 1.578 median index give a strong qualitative test for iron, and fall well in the range of zinnwaldite rather than true lepidolite. The other four lepidolite specimens with high refractive indices are probably between lepidolite and zinnwaldite in composition. Three specimens of book lepidolite had median indices of 1.557, 1.559, and 1.562 and one specimen of curved lepidolite had a median index of 1.560. These values average only a little higher than the purple lepidolite occurring in finegrained aggregates and show that the chemical composition of the different forms vary only to a minor extent. The shape and size of lepidolite gives no clue as to its chemical composition or optical properties. guide noted was color, with the paler and whiter forms having higher refractive indices and less lithia.

Table 18.—Median refractive index (N_{β}) and description of lepidolite from Quartz Creek district

Pegmatite No.	Pegmatite Name	NG	Description
208 215	Opportunity No. 1	1.578 1.565	White (1/4-1/8 inch sheets). White with lilac tinges elongate (3/4 inch) blades
306	Opportunity No. 4	1.557	in a radial aggregate Lilac, 1/4-1/2 inch sheets
306	Opportunity No. 4	1.558	Purple, fine-grained ag- gregates (1/32 inch sheets)
306	Opportunity No. 4	1.564	White with lilac tinge (1/4 inch sheets)
452	Brown Derby No. 1	1.557	Purple, fine-grained ag- gregates (1/16 inch sheets)
452	Brown Derby No. 1	1.555	Purple, fine-grained ag- gregates (1/8 to 1/4 inch sheets)
452	Brown Derby No. 1	1.559	Lilac, book lepidolite (6-inch sheets)
452	Brown Derby No. 1	1.560	Purple, curved lepidolite (1 to 2 inch curved books)
461		1.562	Furple, book lepidolite (2 inch sheets)
535	Brown Derby No. 5	1,560	Silvery white blades (1/4-1/2 inch sheets)
535	Brown Derby No. 5	1.564	White, fine-grained aggregate (1/16 inch sheets)
535	Brown Derby No. 5	1.557	Lilac book lepidolite (5 inch sheets)
637		1.575	White (in aggregate with 1/8-1/4 inch sheets)

Lepidolite is found in amounts that range from a trace to 95 percent.

Only the Brown Derby No. 1 (No. 452) has units containing lepidolite in excess of 10 percent of the rock, the White Spar No. 2 (No. 602) has lepidolite making up 6 to 10 percent of the rock, and all others contained smaller proportions. The units, that contain lepidolite, and are commonly lens shaped, are usually small; several are less than 15 feet long. The Brown Derby No. 1 pegmatite (No. 452) contains by far the largest lepidolite body. This unit is 319 feet long.

Lepidolite was associated with the following minerals: cleavelandite,

quartz, muscovite, perthite, topaz, beryl, microlite, albite, pink and green tourmaline, columbite-tantalite, and apatite. Cleavelandite, the usual form of albite found with lepidolite, is its commonest associate. In two small pods the lepidolite is white and probably zinnwaldite. Topaz, microlite, and colored tourmaline characterize the lepidolite units and are rarely found outside of them.

Lepidolite in many places grows in compact aggregates with cleavelandite or quartz; these minerals appear to have crystallized simultaneously. In other places, lepidolite veins and cuts cleavelandite, quartz, and
perthite. Topaz commonly is surrounded by a rim of lepidolite that may be
in part a product of reaction with the remaining liquid. Lepidolite appears
to have been deposited late in the course of crystallization because in
zoned pegmatites it is confined to the core where it is in part contemporaneous with the quartz and cleavelandite, and in part of later age.

Pyrochlore-microlite

Pyrochlore-microlite is found in 14 pegmatites, all on the east side of Quartz Creek. It does not occur in homogeneous pegmatites, wall zone of zoned pegmatites, or the hanging-wall layer of layered pegmatites; but it is found in the intermediate zone and cores of zoned pegmatites, interior and lower layer of layered pegmatites, and units of pegmatites which change in composition along strike. The most favorable place is the core of zoned pegmatites, as 7 of the 14 pegmatites that contain pyrochlore-microlite occur in this unit. Pyrochlore-microlite commonly occurs in a few scattered crystals, except in pegmatite No. 217 (Opportunity No. 1 claim) and pegmatite No. 452 (Brown Derby No. 1). In pegmatite No. 217 it occurs in concentrations of 10 or 12 crystals in cleavelandite and quartz, whereas

pegmatite No. 452 contains 0.35 percent microlite / in a central lepidolite

unit.

Pyrochlore-microlite is light yellow, light greenish yellow, olive green, light brown, or dark brown. The crystals are from 0.01 to 0.4 inch in diameter. In massive fine-grained lepidolite the crystals are anhedral, but in quartz and cleavelandite they are euhedral with well-developed to distorted octahedrons, o(lll), and modified dodecahedrons, d(ll0).

Pyrochlore is essentially NaCaCb₂O₆F, and microlite is essentially (Na, Ca)₂ Ta₂O₆(O, OH, F). The two species form an isomorphous series with the columbium-rich members called pyrochlore and the tantalum-rich members microlite. Besides the elements given in the above formulas, oxides of some of the following elements may comprise several percent of the mineral: K, Mg, Fe, Mn, Sb, Ce, La, Di, Er, Y, Th, Zr, U, Ti, Sn, and W.

Considerable work has been done on the microlite from the Brown Derby No. 1 (pegmatite No. 452), especially during World War II, when it was mined along with the lepidolite. An analysis of microlite from this pegmatite by J. G. Fairchild has been previously reported / and is given below:

_/ Eckel, E. B., and Lovering, T. S., Work of Eckel, Lovering, Fair-child, Microlite from Ohio City, Colorado: Report of the Committee on the Measurement of Geologic Time, pp. 78-79, 1935.

68.47	Be0	none
4,45	P205	none
	Cão	9.19
	MgO	.04
1.69	PbO	.40
2.40	Na ₂ 0	2.94
20.7		.25
	F	1.51
	Cl	none
.11	Insol. 4 SiO2	.92
none	~	2.84
.04	2	98.41
.07	Less 0 = F	.64
		97.77
	4,45 .10 .95 1.69 2.40 none .10 1.91 .11 none .04	4.45 P205 .10 Ca0 .95 Mg0 1.69 Pb0 2.40 Na ₂ 0 none K ₂ 0 .10 F 1.91 C1 .11 Insol. * Si0 ₂ none H ₂ 0 .04 .07 Less 0 = F

The specific gravity of this material is 5.604, and its index of refraction, as determined by J. J. Glass of the Geological Survey, is close to 1.93. This analysis indicates that the mineral is microlite having a high ratio of tantalum to columbium. As would be suspected from the uranium content this mineral is highly radioactive and can be easily detected with a Geiger-Mueller counter.

The material from the Brown Derby No. 1 pegmatite (No. 452) is dark brown to light brown. An olive-green specimen from the Brown Derby No. 5 pegmatite (No. 535) was analyzed spectrographically, and it was found that the tantalum was more abundant than the columbium. This specimen is also on the microlite side of the series. It gave a positive test with the Geiger-Mueller counter, but not as strong as that given by the dark-brown variety.

One light-greenish-yellow specimen from pegmatite No. 461 gives no reaction to the Geiger-Mueller counter. The dark color of the pyrochlore-microlite may be caused by the radioactivity of the uranium, as are the brown halos surrounding the dark pyrochlore-microlite in the lepidolite.

Similar observations were made by Adams _/ at the Harding mine near

_/ Adams, J. W., Personal communication.

Dixon, N. Mex., where he found that light microlite was not radioactive but dark microlite was.

Pyrochlore-microlite is associated in the Quartz Creek district with: cleavelandite, lepidolite, quartz, and muscovite. The two types of occurrences are: (1) with massive fine-grained lepidolite, and (2) with cleavelandite and smoky quartz. In eight pegmatites it occurs in lepidolite and in six in cleavelandite and quartz. In most places, where it occurs in lepidolite, cleavelandite is present in minor amounts; where it occurs in cleavelandite, lepidolite is a minor constituent. Lepidolite is present in 12 out of the 14 pyrochlore-microlite-bearing units, and cleavelandite in 13. The occurrence of pyrochlore-microlite with either lepidolite, cleavelandite, or both, is common in pegmatites in other areas. It occurs with lepidolite and cleavelandite at the Bob Ingersoll mine, Penfield County, S. Dak., with cleavelandite and lepidolite at the Tin Mountain mine, Custer County, S. Dak., with cleavelandite and a lithium mica in the Harding mine, near Dixon, N. Mex. _/, with cleavelandite at the

_/ Adams, J. W., Personal communication.
Rutherford and Morefield mines, near Amelia, Va/, and with lepidolite
_/ Glass, J. J., The pegmatite minerals from near Amelia, Virginia: Am. Mineralogist, vol. 20, p. 753, 1935.
and albite (type not defined) from a pegmatite at Topsham, Maine _/.
_/ Palache, C., and Gonyer, F. A., Microlite and stibiotantalite from Topsham, Maine: Am. Mineralogist, vol. 25, p. 412, 1940.

Topaz

Topaz is relatively rare in pegmatites and the pegmatites of many districts do not contain this mineral. The Quartz Creek district contains eight topaz-bearing pegmatites, which is less than half of a percent of all pegmatites mapped.

The topaz is milky white, though some has a greenish stain on the outside. The crystals are subhedral to euhedral and are predominantly long tapering prismatic crystals. They are usually 4 to 8 inches in diameter and a foot or two in length, but specimens have been found which attained 12 inches in diameter and 4 feet in length. The prism faces, m(110), and the pyramid faces, i(223), are the best developed. The basal pinacoid, c(001), was noted on a few specimens, and probably other faces could be found with continued study. Basal cleavage is well developed on most specimens. The lower index of refraction (N₂) was determined to range from 1.616 to 1.618 (table 19) on four topaz specimens from three different pegmatites.

Winchell _/ and Pardee, Glass, and Stevens _/ have shown that the indices

J Winchell, A. N., Elements of Optical Mineralogy: Part II, p. 199, 1947.

J Pardee, J. T., Glass, J. J., and Stevens, R. E., Massive low-fluorine topaz from the Brewer mine, South Carolina: Am. Mineralogist, vol. 22, pp. 1063-1064, 1937.

of topaz increase with the increase of water content and decrease of the fluorine content. The indices of topaz in the Quartz Creek district show that there is a uniform water and fluorine content, and by interpolating from a table given by Winchell \int that they contain between 17.0 and 18.5

[/] Winchell, A. N., op. cit., p. 199, 1947.

percent fluorine and 0.9 to 1.5 percent water.

Table 19.-Lower index of refraction (Na) of topaz from the Quartz Creek district

Pegmatite Number	Pegmatite Name	Na
215	Opportunity No. 1	1,616
452 452	Brown Derby No. 1	1.617
452	Brown Derby No. 1	1.618
1574	Bucky	1.616

Topaz is found only in lepidolite-bearing pegmatites and is directly associated with the following minerals in one or more pegmatites: lepidolite, cleavelandite, quartz, muscovite, beryl, perthite, and tourmaline. It is always found with the first three minerals. Purple lepidolite commonly forms a coating around topaz. Its association with lepidolite would be expected as this mineral also contains fluorine in its chemical composition. Stevens _/ made analyses on 17 different lepidolite specimens and

found the fluorine content ranges from 4.09 to 9.19 percent, the average is 7.03 percent. On the Opportunity No. 1 claim, pegmatite No. 215, some of the topaz has a thin pale-green micaceous coating of polylithionite. This specimen of polylithionite is biaxial negative and has N $_{\rm S}$ = 1.558 and N $_{\rm Y}$ = 1.565. Both polylithionite and lepidolite appear to form as a product of reaction between the early-formed topaz crystals and the residual liquid, and corrode the surface of the topaz crystals in some places. The topaz always occurs with the lepidolite, as the fluorine evidently comes off from the original magma at the same time as the lithium. Thus, in a pegmatite district in which both topaz and lithium minerals are present they should be associated. Less than 0.5 percent of the pegmatites of the Quartz Creek

_/ Stevens, R. E., New analyses of lepidolites and their interpretation: Am. Mineralogist, vol. 23, p. 615, 1938.

district contain topaz, yet almost 50 percent of the lepidolite-bearing pegmatites have topaz, and all of the topaz-bearing pegmatites contain lepidolite.

Gahnite

Gahnite, the zinc spinel, is a rare mineral found in only eight pegmatites. It occurs in homogeneous pegmatites, in intermediate zones and cores of zoned pegmatites, and in layered pegmatites.

Gahnite is greenish black to dark green, and occurs in anhedral masses. The crystals are from 0.03 to 0.75 inch in diameter. The mineral is green, isotropic, and has an index of refraction of 1.81 ± 0.005.

Gahnite does not appear to have any favored mineral association. In pegmatite No. 1540 it is found with massive quartz, and muscovite, in pegmatite No. 1574 with albite, quartz, muscovite, and beryl, and in pegmatite No. 452 with albite columbite—tantalite, monazite, tourmaline, garnet, biotite, and quartz. The pegmatites containing this mineral are widely scattered and its presence depends upon the availability of zinc in the pegmatitic liquid.

Gahnite is rare or absent in most pegmatite districts, but has been reported from the Tims Hill deposit in Connecticut _/.

____ Foye, W. G., Mineral localities in the vicinity of Middletown, Connecticut: Am. Mineralogist, vol. 7, p. 9, 1922.

Allanite

Allanite is rare in the Quartz Creek district. In the Black Wonder (No. 847) it occurs in several pods, a few feet thick and about 10 feet long. This pegmatite is over 6,700 feet wide and 12,600 feet long and the

pods represent only a minute fraction of the total pegmatite. The pods are quartz (about 90 percent) and albite. A few scattered crystals of allanite occur in smoky quartz.

The allamite is in prismatic crystals, with a square cross section, as much as 0.5 inch across and 2 inches long. The mineral is black, has a shiny lustre, and is ringed by a reddish-brown decomposition product. At least three different substances are observed under the microscope. One is isotropic, reddish brown in color, and has a refractive index a little less than 1.62; the second is anistropic and colorless; and the third is isotropic and grayish green. These observations are in agreement with those of Hitchens who describes the allamite from Fitchburg, Mass.

____ Hitchens, C. S., The pegmatites of Fitchburg, Massachusetts: Am. Mineralogist, vol. 20, p. 18, 1935.

Unidentified mineral

An unidentified shiny greenish-black mineral, believed to be a new species, was found at the Bucky pegmatite (No. 1574). This mineral occurs in scattered pockets in the mica zone. A total of 17 pounds was collected during the mining operation. This mineral is associated with muscovite, altered feldspars, quartz, monazite, and columbite-tantalite. It appears to be most closely associated and in places intergrown with columbite-tantalite.

The mineral has a conchoidal fracture, and superficially resembles samarskite, fergusonite, or euxenite. A powder X-ray film was compared with the files of known minerals in the X-ray libraries of the Geological Survey and Columbia University with negative results. A spectrographic analysis made by A. T. Meyers in the Geological Survey Trace Elements

Section Washington Laboratory indicates with the order of magnitude of concentration, the following components:

Cb205	XO.O	Y_0_	.X
Cb ₂ 0 ₅ Ta ₂ 0 ₅ U ₃ 0 ₈	X.0	Sno	.X
U308	X.O	Lazoz	Traces
Tho2	,X	La ₂ 0 ₃ Ti0 ₂	· X
MnO	.X	Pb0	.X
Fe ₂ 0 ₃	.X	Sb203	.OX
CaQ.	.X	Sb203 Mg0	Trace
Zro2	X		

Looked for but not found: Na, Bi, W, and P.

A determination of equivalent uranium made by measuring the radiation from uranium and thorium, gave a value of 11 to 12 percent which would show that the X.O given for U₃O₈ would be nearly 10 percent. Like many other radioactive columbo-tantalates this mineral is metamict.

This mineral, in thin section, is pale yellow, isotropic, and has an index of refraction of 1.80 ± 0.05.

The mineral has a specific gravity of 3.8, and some of it contains small cavities lined with a fine-grained yellow material.

Other minerals

Finely-grained chlorite occurs in nine small pegmatites in the Quartz Creek district. It comprises several percent of these small pegmatites.

The grain size is from 0.03 to 0.06 inch. Chlorite is found in fine-grained albite-rich pegmatites and is in part an alteration product of biotite.

Samarskite, or a similar mineral, such as euxenite, is present in seven pegmatites. Only one or two crystals of this mineral, 0.03 to 0.5 inch long, are found in each pegmatite. It has refractive indices above 1.83 and in feldspar is surrounded by a reddish halo. This mineral is commonly associated with smoky quartz, and is strongly radioactive.

Epidote occurs in three pegmatites as fine-green veinlets, and was

introduced into the pegmatite after its solidification.

Light-blue apatite, as crystals 0.25 to 0.5 inch across, occurs in three pegmatites. In most pegmatite districts apatite is very common, but in the Quartz Creek district it is very rare.

Light purple fluorite is found in two pegmatites as grains about 0.06 inch in size. It is extremely rare.

Spodumene and amblygonite occur only in the Bazooka pegmatite (No. 424) in a circular core unit 20 feet in diameter. Spodumene in white lathlike crystals is found on the small dump. No amblygonite was found by the writer, but was observed by Hanley _/.

_/ Hanley, J. B., et.al, op. cit., pp. 66-68, 1950.

Phosphates of the lithiophilite—triphylite series and their alteration products are found in the Bucky pegmatite (No. 1574). Two crystals were noted in the mica zone and one crystal was found in perthite-quartz pegmatite adjacent to a subsidiary core segment, approximately 2,500 feet to the southeast.

Betafite is reported by Hanley _/ associated with monazite, gahnite,

and columbite-tantalite in the Brown Derby No. 1 pegmatite (No. 452). It was not observed by the writer and is probably very rare.

Alteration of wall rocks

The alteration of wall rock by the introduction of pegmatitic materials is common in many districts. Jahns _/ describes impregnation of quartzite

_/ Jahns, R. H., Mica deposits of the Petaca district, Rio Arriba County, New Mexico: N. Mex. Bur. Mines Bull. 25, pp. 52-54, 1946.

and mica schists in the Petaca district by muscovite, microcline, and plagioclase to the extent that the contact between some of the pegmatites and the
country rock is gradational. The formation of muscovite and tourmaline in
the country rock adjacent to the pegmatites in New England has been briefly
mentioned _/. Numerous pegmatites in the Black Hills of South Dakota show

_/ Cameron, E. N., Larrabee, D. M., Page, J. J., Stewart, G. W., and Shainin, V. E., Pegmatite investigations in Maine, New Hampshire, and Connecticut, 1942-1945: U. S. Geol. Survey Prof. Paper. (In preparation.)

abundant alteration at the wall rock-pegmatite contact _/. The Helen

_/ Page, L. R., et al, op. cit, (In preparation).

Beryl pegmatite in Custer County has patches of granulite along its sides from a few inches to 6 feet thick. The granulite varies in composition from place to place and consists of quartz (30 to 70 percent), muscovite (5 to 30 percent), biotite (2 to 15 percent), and minor quantities of tourmaline and apatite. The Elkhorn pegmatite, also in Custer County, has intensely tourmalized the quartz-mica schist on the hanging-wall side of the pegmatite.

In the Quartz Creek district, however, there has been practically no alteration of the country rock adjacent to the pegmatite. The only noticeable alteration was of hornblende gneiss, which appeared slightly more friable adjacent to the contact. The three main types of country rock, hornblende gneiss and tonalite, granite, and quartz monzonite, are equally free of alteration.

Alteration of the country rock might not be expected in the granite and quartz monzonite, as both rocks contain essentially the same minerals as the pegmatites. On the other hand, the hornblende gneiss and tonalite are

markedly different in chemical composition from the pegmatites. Similar hornblende rocks in other districts have been intensely altered. Jahns ______

/ Jahns, R. H., op. cit., p. 54, 1946.

in his description of the Petaca district states, "where amphibole schist lies against pegmatite, as in the Green Peak deposit, it has been converted to a dense aggregate of biotite flakes". McLaughlin _/ in a paper on the

_/ McLaughlin, T. G., Pegmatites of the Bridger Mountains, Wyoming: Am. Mineralogist, vol. 25, p. 53, 1940.

pegmatites of the Bridger Mountains, Wyo., states that all the older pegmatite dikes are accompanied by alteration on the hanging-wall side of the dike, where the percentage of hornblende in the original hornblende schist was greatly reduced and quartz became the predominant mineral.

The type of country rock may affect the kind or amount of alteration, but the type of country rock does not appear to be the prime controlling factor. The only difference between the pegmatites of the Quartz Creek district and those of many other districts that have widespread alteration along pegmatite contacts, is in the composition of the original pegmatite liquid. Tourmaline, apatite, and muscovite are some of the commonest minerals formed in the zone of alteration. Both tourmaline and apatite are among the commoner minerals in most pegmatite districts and may occur in almost every pegmatite; but in the Quartz Creek district tourmaline is a minor constituent of 48 out of 1,803 pegmatites and apatite is found in two. Tourmaline is the only boron mineral found; but phosphorous occurs in apatite, in lithiophilite—triphylite in one pegmatite, in amblygonite in another, and in monazite in twenty—three. Muscovite, though common in the number of occurrences, is small in amount as compared to many other areas.

These facts indicate that the original pegmatitic magma contained little B,

P, and OH, and possibly other volatiles. Alteration of the wall rock is

dependent upon solutions derived from the cooling pegmatite; therefore,

this deficiency of water and other volatiles is a determining factor in

adding new minerals to wall rocks.

The lack of alteration in the Quartz Creek district appears to be caused by the insufficient concentrations of the elements needed to form alteration minerals, though they were available in the pegmatite magma to form rare minerals in the pegmatite during all stages of crystallization.

Distribution of minerals

Some pegmatite districts are important as a source of lepidolite, sheet mica, columbite-tantalite, beryl, or other pegmatite minerals. The granitic pegmatites of most districts consist essentially of perthite, plagioclase, quartz, and muscovite, but not all pegmatite districts have the same assemblage of minor minerals. It is this assemblage of minor minerals and variants of common minerals, such as curved muscovite and colored tourmaline, that indicate the differences in the overall composition in the original source magmas of each district.

Not only do the less common minerals vary from district to district, but from pegmatite to pegmatite. During World War II the U. S. Geological Survey was studying pegmatites that produced critically needed materials and pegmatites were grouped simply according to minerals of economic interest. It was recognized by many investigators that a certain type of mineral would occur in certain groups of pegmatites, that is, the lithium-bearing pegmatites of a district would not be scattered haphazardly throughout an area but would occur in a cluster or groups throughout a district. This grouping of mineralogically similar pegmatites is illustrated in the Black Hills of South Dakota where such well-known lithium producing pegmatites as the

Etta, Peerless, Hugo, and Edison, are all in one group; and the Helen Beryl, Helen Beryl No. 2, and Tin Mountain, in another.

The areal mapping on which this report is based afforded an excellent opportunity to study the distribution of minerals in a medium-sized pegmatite district. A series of maps, (figs. 16, 17, 18, 19, and 20), show the distribution of beryl. tourmaline, curved muscovite, biotite, magnetite, monazite, columbite-tantalite, cleavelandite, topaz, lepidolite, and microlite throughout the Quartz Creek district. Some minerals such as flat muscovite, and garnet are too widespread to be significant, whereas others like chlorite, amblygonite, and spodumene, are too rare to be of use statistically. Each pegmatite that contains at least one crystal of a particular mineral is indicated on the map as a bearer of that mineral. This scheme of representation has the serious defect in that large pegmatites appear to have a greater quantity of the mineral than do the smaller pegmatites. For example, the Black Wonder pegmatite (12,600 feet long by 6,700 feet wide), contains only a few crystals of beryl in two small pockets, yet it appears on the map to be a large beryl-bearing area. facts are emphasized by these maps: (1) the relation of distribution of certain minerals to all the pegmatites in the district, and (2) the constant association of two or more minerals. The associated minerals are grouped on the same figure.

Two hundred and thirty-two beryl-bearing pegmatites are shown in figure 16. The group of beryl-bearing pegmatites in the northwestern corner of the area mapped has a northeastwardly trend; a second group extends northeast from near Opportunity claims in the southwest part of the area mapped; a third group extends from Wood Gulch northwest to the Brown Derby mine; a fourth small group occurs in the vicinity of the Buckhorn mine on

the eastern edge of the district; and a fifth small group is around the Bucky mine in the northeast corner of the district. There are scattered beryl-bearing pegmatites, including the Black Wonder, in the north-central part of the area mapped, which actually contain a little teryl.

Figure 17 shows the location of the groups of tourmaline- and curved muscovite-bearing pegmatites. Tourmaline is most abundant in an area near the Opportunity claims on the southwestern edge of the district, in the vicinity of the Brown Derby No. 1 pegmatite, and adjacent to the quartz monzonite in a small area along the northern edge of the map. Curved muscovite is found only in the northwestern slope of Wood Gulch. It is surprising that these relatively rare minerals are grouped so closely in the hundreds of pegmatites mapped.

Magnetite and biotite are found in 422 and 357 pegmatites, respectively. Almost every pegmatite in the northern part of the area (fig. 18) contains these two minerals, but in the southern part these minerals appear only in small clusters of pegmatites.

Figure 19 shows the distribution of columbite-tantalite-bearing and monazite-bearing pegmatites. Although they are not as abundant as the other minerals studied, the special association of pegmatites containing these minerals are clearly the same. Only one or two crystals of these minerals were formed in each pegmatite. They tend to occur in clusters of several pegmatites, as is apparent from the maps.

Figure 20 shows the grouping of lepidolite-, cleavelandite-, topaz-, and microlite-bearing pegmatites. The intimate association of these four minerals is clear. These minerals are abundant in four main groups of pegmatites: (1) in southwestern part of the district around the Opportunity claims, (2) in the vicinity of the Brown Derby mine, (3) in the

vicinity of the White Spar claims, and (4) in the vicinity of the Bucky property.

Relationship of the pegmatites to the country rock

One of the problems of pegmatites, as with many other igneous rocks, is the effect the country rock has on their occurrence and mineral composition. The pegmatites in the Quartz Creek district occur more commonly in hornblende gneiss and tonalite, and less frequently in granite and quartz monzonite. If the pegmatites were derived from the same magmas as either the granite or quartz monzonite, they would tend to occur along the outer edges of the parent igneous rock and in the adjacent country rock. In the Quartz Creek district, however, pegmatites are found on the edges of both the granite and the quartz monzonite. Also the pegmatites would not penetrate far into the granitic rocks, if they came from the outside.

An equally logical reason for the pegmatites distribution in this area is that the granite and quartz monzonite are less fractured and afford less easy passage to the pegmatite solutions. The pegmatites find zones of weakness to intrude only on the edges of these bodies, where cooling fractures are common. This does not preclude that the pegmatites were not derived from the original magma of either of these two rocks. It does seem, however, that the presence or absence of pegmatites in a particular rock or part of a rock may be the result of the ease of intrusion rather than the source from which they are derived.

A statistical count of the types of pegmatite minerals found in the various rock units was made on the pegmatites in the chief rock types.

All pegmatites that occurred in two or more types of country rock were excluded. Certain minerals such as plagioclase, quartz, and perthite are omni-

present and thus show no differences; whereas others like topaz, gahnite, and spodumene are too rare to give a significant statistical count.

Another problem is the distribution of minerals in groups which may appear to show concentrations of a mineral in one rock type, but which are related to the areal distribution rather than to the rock type. Some minerals such as lepidolite and cleavelandite are not found in pegmatites in either the granite or the quartz monzonite. The small number of pegmatites in which they are found, however, precludes any statement as to the effects of the country rock. Tourmaline, on the other hand, is fairly uniformly distributed in the three dominant types of country rock considering that it is found in only 48 pegmatites. The ratio of tourmaline-bearing to non-tourmaline-bearing pegmatites is 1:30 for hornblende gneiss and tonalite, of 1:21 for granite, and of 1:85 for quartz monzonite.

Beryl-bearing pegmatites show a marked preference for certain types of wall rock. Beryl occurs in 232 pegmatites and though irregularly distributed, it is believed, at least in the case of granite, which is also irregularly distributed in many areas of beryl-bearing pegmatites, that the distribution is wide enough to discount the general areal pattern of beryl occurrences. The ratio of beryl-bearing pegmatites to non-beryl-bearing pegmatites is 1:6.4 in hornblende gneiss and tonalite, 1:20 in granite, and 1:189 in quartz monzonite. These figures show wide variance and suggest that the concentration of the beryl in pegmatites is influenced by the country rock. It was thought that the granite might absorb BeO from the pegmatite liquid, and, therefore, bulk samples were taken from a graphic granite-rich pegmatite (No. 512) and the adjoining coarse-grained granite. The pegmatite sample consisted of perthite (62 percent), quartz (20 percent), albite (15 percent), and muscovite (3 percent). A little more than half of this rock

was graphic granite. The granite was estimated to contain microcline (67 percent), albite (20 percent), quartz (8 percent), and biotite (5 percent). The pegmatite contained a trace of BeO and the granite contained an amount less than was detectable spectrographically (under 0,0001 percent). Bulk samples were taken also of the footwall layer of the Brown Derby No. 1 pegmatite (No. 452) and of the hornblende gneiss, within 8 inches of the contact. Samples were split down from about 1,000 pounds of original material. The pegmatite was estimated to consist of albite (89 percent), quartz (10 percent), tourmaline (1-2 percent), muscovite (less than 1 percent), and garnet (trace). It had an average grain size of 0.12 inch and there was no visible beryl. The analysis showed this pegmatite to contain 0.030 percent BeO and the hornblende gneiss a trace. The results of these two sets of samples should be supplemented by much more data. Graphic granite pegmatites are commonly lean in beryl and the Brown Derby pegmatite had beryl in other units than the one sampled. This small amount of work seems to indicate that granite does not absorb BeO from the pegmatite fluid and thereby cause the difference between the BeO content of pegmatites in granites and in hornblende gneiss. The possibility of the BeO being derived from the country rock has not been investigated. Samples of the hornblende gneiss away from the pegmatite are needed to see if it too contains a trace of BeO. More probably, however, the trace of BeO in the hornblende gneiss is derived from the pegmatite, as BeO tends to be concentrated in the last stages of magmatic differentiation _ and is chiefly

_/ Rankama, K., and Sahama, T. G., Geochemistry, p. 443, 1950.

found throughout the world in granites and nepheline syenites.

Origin

The problem of the origin of pegmatites is complex and involves not only the method by which they are derived from the original magma, but also their crystallization. Pegmatites appear in many areas to be related areally to large bodies of intrusive rock. Most of these intrusive rocks crystallized from magmas of silicic composition and thus pegmatites are commonly associated with igneous rocks such as granites. Goranson _/ has

shown that certain natural rhyolitic glasses may contain 8 to 10 percent water. Thus, from a granitic magma, containing only 1 to 2 percent water, the amounts of water released in the later stages by slow crystallization of such a magma would be very large. In addition to water, the other volatile elements such as F, Cl, B, and P, would be concentrated. The alkalies, Na, K, Li, Cs, and Rb, also tend to be concentrated in the later stages of crystallization.

Granitic pegmatites have some of the properties of granites and some of various types of veins. They appear to be an intermediate type and have been correlated by various writers to both igneous rocks and veins. Beryl, a typical pegmatite mineral, is found in both granites and quartz veins as well as in pegmatites suggesting a continuous gradation between these rock types. Beryl, for example, is found in the granites on Mt. Antero, Colo.

_/ Adams, J. W., Beryllium deposits of the Mt. Antero region, Chaffee County, Colorado: U. S. Geol. Survey Bull. (In preparation.)

and in the Victorio Mountains, N. Mex. _ and beryl-quartz veins are found

Holser, W. T., Unpublished Geol. Survey report on the occurrence of helvite and beryl in the Victorio Mountains, New Mexico.

in the Victorio Mountains \mathcal{J} , the California vein on Mt. Antero \mathcal{J} , the

Holser, W. T., op. cit.
Adams, J. W., op. cit.

Boreana vein, Arizona J, and Kazakhstan, Russian J.

/ Hobbs, S. W., Tungsten deposits in the Boreana district and the Aquarius Range, Mohave County, Arisona: U. S. Geol. Survey Bull. 940-I, p. 254, 1944.

Most of the pegmatites in the Quartz Creek district are simple pegmatites composed of minerals typical of granites—perthite, quartz, plagioclase, muscovite, and garnet. The unusual "distinctive" minerals of pegmatites rarely are found. Such minerals as cleavelandite, lepidolite, topaz, microlite, gahnite, and columbite—tantalite occur in less than 2 percent of the pegmatites in the Quartz Creek district. The predominant mineralogical difference between pegmatites in many districts and their associated granites is that the pegmatites are somewhat higher in muscovite, indicating a higher water content of the original liquid. In the Quartz Creek district, however, muscovite is a relatively minor mineral, suggesting that the original magma was water—poor. The few pegmatites that have minerals containing other volatile elements are the result of a later stage of segregation and crystallization,

More than 90 percent of the pegmatites in the Quartz Creek district have an average grain size of less than 1 inch. This texture is commonly aplitic, and may resemble that of a fine- to coarse-grained granite. Many of the pegmatites in the Quartz Creek district resemble typical igneous rocks and have been mapped by Crawford and Worcester \int as granite in the

_/ Crawford, 'R. D., and Worcester, P. G., Geology and ore deposits of the Gold Brick district, Colorado: Colorado Geol. Survey Bull. 10, 1916.

northern part of the area. The lepidolite-bearing Brown Derby pegmatites, on the other hand, are much coarser textured and contain many minerals distinctive of pegmatites that have been described as showing "abundant evidence of hydrothermal replacement," by Landes _/.

_/ Landes, K. K., Colorado pegmatites: Am. Mineralogist, vol. 20, p. 333, 1935.

The pegmatites of every district have distinctive characteristics that usually are reflected in the rare minerals. For example, the pegmatites of the Black Hills are in general rich in tourmaline, muscovite, apatite, and other phosphate minerals. Very little topaz and lepidolite are found in this large district. These minerals show that the original magma was comparatively rich in P, B, OH, and poor in F. The pegmatites of the Quartz Creek district, on the other hand, differ from those in the South Dakota district and many others in that they are relatively lean in muscovite, biotite, tourmaline, and phosphate minerals, and are relatively rich in topaz, lepidolite, and columbite-tantalite. These minerals indicate that the original magma of the Quartz Creek district pegmatites was lean in OH, B, and P, and comparatively rich in F, Ta, and Cb. In the entire Quartz Creek district, only three pegmatites contain enough muscovite to be considered as sources of scrap mica; the content is commonly 0.5 to 3 percent. Biotite is less than I percent. The lack of these two minerals in most pegmatites indicates that the pegmatites of this area contain relatively Tourmaline is found in only 48 pegmatites and the content is only a fraction of 1 percent. The dominant phosphorous mineral is monazite, which is present only as a few small crystals in 24 pegmatites. Other phosphate minerals are apatite, amblygonite, and lithiophilite—triphylite, which are found in only one or two pegmatites and are exceedingly rare. Fluorine is a constituent of topaz in 8 pegmatites, lepidolite in 17 pegmatites, microlite in 14 pegmatites, and in fluorite in 2 pegmatites. Columbium and tantalum are present in columbite—tantalite in 29 pegmatites; in microlite in 14 pegmatites, and in samarskite (?) in 7 pegmatites, and in an unidentified mineral in 1 pegmatite.

Within a particular district considerable variation is shown in the areal distribution of pegmatite minerals. As previously discussed, pegmatites of some of the rare elements occur in groups. Thus, beryl-bearing pegmatites are found in several clusters scattered over the area (fig. 16). This is also true for some of the other rare minerals as lepidolite, topaz, cleavelandite, microlite, columbite-tantalite, monazite, tourmaline, and curved muscovite (figs. 17, 18, 19, and 20). Pegmatite magma evidently escapes from a particular part of the chamber of its parent granitic magma in a specific direction dividing into separate units before final emplacement. This would account for the distribution of certain types of pegmatite minerals in one area. The distribution of these minerals in different groups is related to their origin. As the parent granite magma cools, pegmatite liquids are segregated in different parts of the magma and escape from various parts at different times. The pegmatite magma which is driven from the parent chamber earliest probably contains less volatile material, and forms the greater part of all pegmatites-those that most closely resemble other granitic rocks. The minerals found in any pegmatite depend on the original composition of the material segregated in a pocket, and on the stage of crystallization at which it was derived. The more highly volatile constituents are in the later derivations and form the few rare

pegmatites. Many minerals are almost always found in close association with another mineral because these minerals may contain common ions which make their association imperative, or elements that are concentrated at the same stage in crystallization. Minerals containing common ions include: lepidolite, topaz, and microlite—all of which contain F; and lepidolite, and colored tourmaline which contain Li. Minerals which owe their association to elements segregated at approximately the same stage are columbite—tantalite and monazite. In places where these two minerals are not associated it probably reflects a lack of elements to form one or the other mineral.

The commoner minerals are segregated continuously or recurringly through the differentiation of the parent granite. In pegmatites that contain the rarer elements, minerals such as plagioclase, perthite, quartz, and muscovite are still the predominant minerals. The rarer elements such as Li, Cs, Rb, F, Cl, Cb, and Ta are probably in the more soluble part of the pegmatite material and certainly are among the last to crystallize.

The reaction of these late crystallizing rest solutions on the earlier crystallized material causes embayment and veining of the earlier crystallized minerals. This is given as proof of replacement by later solutions by some authors _____. Whether the first magmatic part of a pegmatite was re-_____.

[,] Mineral replacement in pegmatites: Am. Mineralogist, vol. 12, pp. 59-63, 1927.

Landes, K. K., Paragenesis of the granite pegmatites of central Maine: Am. Mineralogist, vol. 10, pp. 355-411, 1925.

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Jahns, R. H., Mica deposits of the Petaca district, Rio Arriba County, New Mexico: New Mexico Bur. Mines Bull. 25, pp. 72-75, 1946.

placed by hydrothermal solutions brought in after the first part of the pegmatite had solidified, or whether the earlier crystallized minerals were acted on by a rest solution that became unstable with the earlier crystallized minerals as the pressure and temperature changed the results would give identical textural relations with the earlier formed minerals being corroded and veined by the later. Evidence on whether pegmatites were formed in a closed system or by a series of hydrothermal replacements must depend on the mineralogical and structural relations rather than on a textural study. Studies of internal structure of the pegmatites and the order of sequence of mineral groups are markedly uniform and can be correlated from district to district as well as among pegmatites in the same district. The uniformity of sequence of mineral assemblages, and the relationship of the minerals themselves point to a similarity of conditions which would not be expected from hydrothermal replacement. A discussion of hydrothermal replacement versus forming in situ has been discussed thoroughly by Cameron, Jahns, McNair, and Page ___; Stockwell ____; Hanley _____; and Page _______, and will

Hanley, J. B., et al., op. cit., pp. 7-9, 1950.

Page, L. R., et al., op. cit. (In preparation.)

not be repeated here.

Reserves

Reserves of pegmatite minerals are difficult to estimate because normal procedures of sampling can not be used. The grade, however, can be obtained by measuring the areas of industrial minerals exposed on pegmatite surfaces and relating it to the total exposed area. The percent of mineral exposed may be converted to a weight percent of mineral by making proper corrections for specific gravity.

The tonnage of rock containing an industrial mineral can be calculated from a detailed map of the internal structure of pegmatite.

Reserve calculations have been made for all of the industrial minerals in pegmatites of the Quartz Creek district. The minerals that sustained mining operations in the past were lepidolite or scrap muscovite and beryl. Potash feldspar might sustain mining operations, if transportation were less costly.

The total reserves of clean hand-cobbable feldspar are estimated to be 795,600 tons of potash feldspar and 9,740 tons of soda feldspar. These feldspar reserves are all in pegmatite units containing more than 25 percent feldspar, in grains greater than 12 inches in length. The minimum size of a unit included in these calculations was 200 feet long and 40 feet thick.

In addition, there is considerable feldspar recoverable by milling. Most of this feldspar is in the form of graphic granite. A number of these pegmatites pose considerable transportation difficulties because they are several miles over mountainous terrain from the nearest road. A total of 251,300,000 tons of milling grade feldspar are calculated in 40 pegmatites. This is an average of 6,028,000 tons per pegmatite. The largest tonnage is in the Black Wonder (No. 847), which has 225,200,000 tons of milling

feldspar,

There is no sheet mica and very little scrap mica in the Quartz Creek district. The total reserves of scrap muscovite is estimated to be 13,500 tons of which 1,400 tons is recoverable by hand methods. The scrap mica obtained is a byproduct of the mining of beryl-bearing pegmatites. There are only three pegmatites (the Bucky, the Buckhorn, and the Beryl and Rare Minerals Lode) which contain either enough muscovite or muscovite in large enough pieces to be considered recoverable as a byproduct of beryl mining. The muscovite reserves of these three pegmatites are calculated as 1,400 tons.

The total beryl reserve estimated for the Quartz Creek district is about 350 tons. The second largest pegmatite in the district was estimated to contain 160 tons of beryl as of November 1949. This deposit is twice as large as the reserves of any other pegmatite. Pegmatite No. 452 (Brown Derby No. 1 dike) contains 75 tons, the second largest beryl reserve in the area and pegmatite No. 538 is third with 40 tons of beryl. Thirty-eight of the pegmatites contain less than 10 tons of beryl. Of the total, 350 tons of beryl, probably 325 tons are hand cobbable.

Some of the pegmatites that have beryl reserves also contain lepidolite, microlite, topaz, columbite-tantalite, and monazite. The lepidolite reserves of the entire area amount to 3,560 tons. The largest deposit is the Brown Derby No. 1 dike (pegmatite No. 452) which has 1,600 tons of reserve. Only four lepidolite deposits have reserves over 100 tons. The topaz reserves are 900 tons, those of microlite 900 pounds, columbite-tantalite 4,000 pounds, and monazite 400 pounds. The reserves of the last three minerals, because they occur in such small widely distributed quantities, are very difficult to calculate.

Prospecting for beryl

Of the 1,803 pegmatites found in the Quartz Creek district, 232 of them contain some beryl. Only one or two crystals of beryl, 0.10 to 0.25 inch in diameter, occur in most of these pegmatites. Of the 232 pegmatites, 42 have more than 2 square inches of beryl exposed, and of these only a very few could be considered as possible sources of appreciable beryl production.

The finding of pegmatites that contain beryl in sufficient quantities to be of commercial value is very difficult. The beryl in this district is commonly white and approximates the physical appearance of feldspar or quartz. It is commonly overlooked by prospectors. Favorable beryl-bearing zones occur in covered areas in many places and diligent prospecting in these favorable areas might uncover worthwhile deposits.

Several broad statements can be made concerning the favorable and unfavorable areas for finding beryl in pegmatites in the Quartz Creek district. Granite and quartz monzonite appear definitely unfavorable as a host rock for beryl pegmatites. Only three pegmatites were found in granite that contained beryl and none of the exposures contained as much as 2 square inches of beryl. Only one pegmatite in quartz monzonite contained beryl. All pegmatites in the Quartz Creek district that contained as much as 2 square inches of beryl had hornblende gneiss wall rocks. The hornblende gneiss covers a large area that is favorable for detailed prospecting.

Detailed studies also indicate that beryl is most common in albiterich pegmatites. The albiterich pegmatites are almost universally fine grained. The grain size of the beryl in these pegmatites ranges from 0.06 to 0.5 inch and no beryl has been recovered commercially from them. Perthiterich pegmatites, though less common as a source of beryl, have a large

enough grain size so that most of the beryl can be hand cobbed. The small grain size of the beryl-bearing albite-rich pegmatites precludes economic recovery of beryl under present technological conditions, though at some future date, such beryl-rich rock, if found in sufficient quantities, may be of economic importance and warrant the erection of a mill. Deposits of this type, therefore, can not be expected to be of immediate concern to the prospector.

Some pegmatites have both albite—and perthite—rich parts. The perthite—rich parts may be a source of beryl at the present time but the albite—rich parts, though equally rich in beryl, will be too fine grained to mine. The beryl in pegmatites containing cleavelandite is coarse grained and may be recovered by hand methods, consequently cleavelandite—bearing pegmatites should be prospected carefully.

Graphic granite pegmatites, one of the most common types of perthiterich pegmatites in the district, do not contain beryl and should be avoided in searching for beryl deposits. Beryl favors blocky perthite-quartz pegmatite units or perthite-muscovite-rich units, though the very irregular distribution of beryl may cause one part of a unit to be completely barren of beryl, whereas in other parts it may be abundant. Prospecting for beryl should be most intense in perthite-quartz units, and the edge of quartz cores especially where a feldspar-rich muscovite-bearing unit may be covered by float. Much beryl is found in lepidolite-bearing pegmatites and lepidolite is considered one of the more favorable indicators of beryl in prospecting.

In general, the Quartz Creek district is not too favorable a district to prospect for beryl, because the good mica and feldspar deposits are of small size and production of beryl in the past has been as a byproduct of mining for one or the other of these minerals. There are no sheet mica deposits in the district, the scrap mica deposits are not large, feldspar in large quantities must be recovered by milling of graphic granite, and deposits of hand cobbable feldspar are small. The size of the deposits of these minerals added to the high cost of transportation almost preclude profitable mining operations in this district at 1951 prices. Little feldspar has been sold by local producers, and the high cost of transportation and the small size of the perthite units are not encouraging for the production of feldspar. Because of these factors that concern muscovite and feldspar, the future of this district as a source of beryl is not too bright.

DESCRIPTIONS OF INDIVIDUAL DEPOSITS

Opportunity No. 1 claim (pegmatite No. 215)

The Opportunity No. 1 claim is in the NEL sec. 17, T. 49 N., R. 3 E., New Mexico principal meridian. This prospect is claimed by Earl A. Serry, and is south of the Doyleville road on the eastern slope of a north-trending ridge. The claim covers more than 10 pegmatites, but the main workings are on the largest one, No. 215. This pegmatite is exposed by five small cuts, the largest of which is 37 feet long, 10 feet wide, and 10 feet deep; the other four are each about 8 feet long, 4 feet wide, and 5 feet deep. No minerals have been produced from this prospect.

The pegmatite is a lenticular-branching pegmatite 730 feet in maximum length and 40 feet in maximum thickness. Much of the pegmatite, however, has a thickness of less than 5 feet. It dips 45° to 80° SE., and cuts both fine-grained granite and hornblende gneiss.

The pegmatite is divided into four units along the strike of the body (fig. 21). Albite-quartz-perthite pegmatite forms a unit in the stringers

extending from the northeast, the southwest, and the central part of the pegmatite. This unit has an average grain size of 0.25 inch and is estimated to consist of albite (65 percent), quartz (20 percent), perthite (15 percent), muscovite (less than 1 percent), and one small--0.25 inch--palegreen beryl crystal. Near the south end of the pegmatite, the body widens considerably at the junction of a small northward-trending branch with the main body of the pegmatite. This wide area contains several units. Along the western side is a unit of quartz-albite pegmatite that has an average grain size of 3 to 4 inches, and is estimated to contain quartz (75 percent), albite (20 percent), muscovite (3 percent), perthite (2 percent), and tourmaline (less than I percent). The eastern side of this bulge area is dominantly perthite-cleavelandite-quartz pegmatite, has an average grain size of about 2 to 3 inches, and is estimated to consist of pink perthite (60 percent), white cleavelandite (20 percent), quartz (20 percent), muscovite (less than 1 percent), black tourmaline (less than 1 percent), and garnet (trace). The perthite occurs in crystals as much as 12 inches long. The central part of this bulge, and the remainder of the pegmatite to the north, is cleavelandite-quartz pegmatite. This pegmatite has an average grain size of 1.5 inches and is estimated to consist of white cleavelandite (69 percent), quartz (20 percent), perthite (8 percent). muscovite (2 percent), black and dark green tourmaline (1 percent), lepidolite (less than 1 percent), topaz (less than 0.5 percent), beryl (0.1 percent), and garnet (trace), microlite (trace), columbite-tantalite (trace), and monazite (trace), Locally clots of varying lepidolite content are present. In one 3-foot area lepidolite constitutes as much as 5 percent of the rock. It occurs as fine-grained aggregates and in larger plates 0.25 to 1 inch in diameter. The topaz is milky white, has a good

cleavage and commonly occurs with white beryl. Beryl and topaz have an irregular distribution. Microlite is found in the southern part of this unit with smoky quartz and platy cleavelandite as clusters of distorted of tahedrons, modified by dodecahedrons as much as 0.25 inch in diameter. Probably the microlite content of this small part of the pegmatite is a few pounds per ton. Columbite—tantalite is less common than the microlite, and only a few small crystals, the largest an inch across, were found. Monazite was noted in four crystals and is associated with columbite—tantalite.

Pegmatite No. 417

Pegmatite No. 417 (pl. II) is on the western side of the western branch of Wood Gulch, near the western edge of sec. 2, T. 49 N., R. 3 E., New Mexico principal meridian. It is a few feet above the canyon bottom and can be reached by the secondary road running up Wood Gulch. It is opened by one small pit 18 feet long, 4 feet wide, and 4 feet deep at its western face.

The pegmatite is irregular, and is approximately 370 feet long and 105 feet in maximum width. Part of the irregularity is the result of its exposure on a dip slope. The pit, which is near the center of the pegmatite, cuts completely through the body, and at this point the pegmatite is 6 feet thick. The pegmatite intrudes hornblende gneiss, and has a wall zone and a core. The wall zone is the greater part of the body and has an average grain size of 0.12 inch. It consists of albite (55 percent), quartz (35 percent), perthite (5 percent), muscovite (5 percent), and beryl (0.75 inch crystal). The core has an average grain size of 3 inches and consists of perthite (45 percent), albite (42 percent), quartz (8 percent),

muscovite (5 percent), beryl (0.2 percent), and one small piece of samar-skite. The beryl forms pale green, euhedral crystals that range in size from 0.25 inch by 0.25 inch to 2 inches by 4 inches. A grain count of 101 square feet of the core contained 0.2 percent beryl in 33 crystals. The beryl reserves of this unit are calculated as 3.8 tons.

Brown Derby Dike No. 1 (pegmatite No. 452)

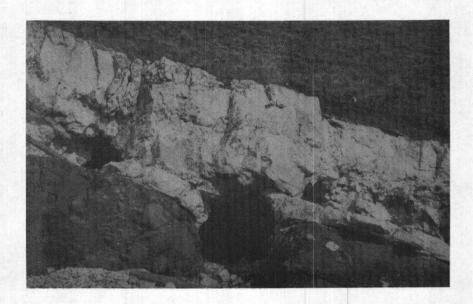
The Brown Derby Dike No. 1 is the easternmost of the Brown Derby group of dikes, in sec. 3, T. 49 N., R. 3 E., New Mexico principal meridian. This dike is a few hundred feet below the crest of a long ridge at 9,300 feet on the east side of Quartz Creek. An access road was built by the Federal Government from Colorado State Highway 1622 to the workings, a distance of approximately 2 miles.

The pegmatite has been described by J. B. Hanley _/. Since this mapping

In May and June, 1950, the U. S. Bureau of Mines drilled two core holes. Diamond-drill hole No. 1 (fig. 22) is inclined 60 degrees, has a bearing of N. 77° W., and was drilled from the dump of dike No. 1. The hole has a length of 208.5 feet, and cut both dikes No. 2 and No. 3. Diamond-drill

_/ Hanley, J. B., Heinrich, E. W., and Page, L. R., Pegmatite investigations in Colorado, Wyoming, and Utah, 1942-1944: U. S. Geof. Survey Prof. Paper 227, pp. 68-71, 1950.

was finished (December 1944), the high-grade pod-shaped concentration of lepidolite in the lepidolite-quartz-cleavelandite core has been completely mined out (pl. XI). In addition, the lepidolite-quartz-cleavelandite core in tunnel 2 has been mined approximately 30 feet east of the former face to a width of 20 feet. This additional work was completed by the Hayden Mining Company before it ceased mining in 1945.



Mouth of stope of the Brown Derby No. 1 pegmatite (No. 452) showing contact of pegmatite with hornblende gneiss.

hole No. 2 was drilled from the hillside above tunnel 3 and cut dikes Nos. 1, 2, and 3. The bearing of the hole is N. 88° W., and the inclination is a minus 65 degrees.

The Brown Derby No. 1 dike is a lenticular and branching pegmatite, with two branches at its southern end. It is exposed for a total length of 913 feet.

The west branch is exposed in a series of six prospect pits. It is made up of a wall zone of albite-quartz pegmatite and a core of albite-quartz-biotite pegmatite.

The albite-quartz pegmatite forms nearly all of the west branch of the dike. It has a grain size of 4 to 6 inches and is estimated to contain quartz (52 percent), cleavelandite (45 percent), muscovite (3 percent), and less than 1 percent of garnet, tourmaline, and lepidolite.

The albite-quartz-biotite pegmatite of the core is exposed in pit No.

11, and is 20 feet long by 1.5 feet wide. The grain size of the core ranges

from 0.12 to 0.25 inch. The unit contains albite (86 percent), quartz (5

percent), biotite (4 percent), monzonite (2.2 percent), columbite-tantalite

(1.4 percent), gahnite, the zinc spinel (1 percent), and less than 1 percent garnet, tourmaline, and trace quantities of fluorite and betafite.

Monazite forms well developed euhedral crystals, 0.25 to 1.5 inches across,

and columbite-tantalite are tabular crystals, 0.25 to 1 inch across.

The main part of the pegmatite is composed of six different units, identified from hanging wall to footwall as follows: perthite-albite-quartz pegmatite (hanging-wall layer), hanging-wall quartz pod, curved lepidolite layer, lepidolite-microlite pod, quartz-cleavelandite-lepidolite-topaz layer, and albite pegmatite (footwall layer).

The perthite-albite-quartz pegmatite makes up the entire width of the

northern part of the dike and the eastern branch of the southern part of the dike. Elsewhere in the lepidolite-bearing part of the dike, it occurs as the hanging-wall layer, with the exception of the vicinity of the inclined shaft where it is missing. At this point the quartz pod takes the place of the perthite-albite-quartz pegmatite. This pegmatite has an average thickness of about 8 feet. In the lepidolite-bearing part of the pegmatite it does not exceed 4 feet. The grain size of the unit is about 12 inches, and the composition is estimated to be perthite (40 percent), albite (30 percent), quartz (20 percent), muscovite (10 percent), beryl (0:1 percent), and trace of tourmaline. Beryl crystals range from 0.25 to 2 inches in diameter.

The hanging-wall quartz pod is approximately 84 feet long and 2 feet thick. Its grain size is 24 inches, and the constituents are estimated as quartz (70 percent), cleavelandite (25 percent), and lepidolite (5 percent). The lepidolite is in sheets from 3 to 4 inches across.

The curved lepidolite layer is the uppermost of the lepidolite-bearing units; it is 190 feet long and averages 2 feet thick. This unit forms the back of the extension of tunnel 2 and much of the back of the stoped area in the inclined shaft. The grain size of the curved lepidolite layer is 4 to 5 inches and the composition is cleavelandite (44 percent), quartz (40 percent), curved lepidolite (15 percent), topaz (1 percent), less than one percent muscovite and tourmaline, and a trace of apatite. The curved lepidolite ranges from 0.25 to 2 inches across, and the topaz ranges from 4 to 8 inches across.

Two principal lepidolite-microlite pods are known. In addition, smaller pods have been exposed in pits and trenches. The largest pod was mined in the inclined shaft and the other was discovered underground and

mined by tunnel 2. The pod at the inclined shaft was approximately 60 feet long with a maximum width of 8 feet. It was mined from the inclined shaft for a total length of about 170 feet down the dip. The pod exposed in tunnel 2 is approximately 30 feet wide by $6\frac{1}{2}$ feet thick, and is present in in the face of the tunnel. The mined length is approximately 40 feet.

The average grain size of the unit is 1 inch, and the average composition is cleavelandite (43 percent), lepidolite (40 percent), quartz (15 percent), topaz (2 percent), microlite (0.35 percent), and a trace of beryl. The lepidolite is in crystals 0.03 to 0.12 inch across, and is irregularly distributed within the pods. Microlite crystals range from less than 0.01 inch to 0.25 inch in diameter and are in shoots within the pods. This unit in tunnel 2 is extremely low in microlite, and has also a low lepidolite content. The pod has been nearly mined out in the incline; the remainder shown in the bottom is pinching to the southeast.

The quartz-cleavelandite-lepidolite-topaz layer is the footwall part of the lepidolite-bearing units of the pegmatite. It is exposed on the surface for 319 feet and has an average thickness of 2 feet. The unit has an average grain size of 4 to 6 inches and consists of quartz (55 percent), cleavelandite (25 percent), lepidolite (10 percent), topaz (10 percent), less than 1 percent muscovite, and less than 0.1 percent beryl. Lepidolite is in flat books ranging from 1 to 7 inches across and topaz crystals are as much as 42 inches long. Beryl is in crystals ranging from 1 to 4 inches in diameter.

The albite pegmatite (footwall layer) has an average thickness of 1.5 feet. It occurs discontinuously along the lepidolite-bearing part of the pegmatite. It has a grain size of 0.25 inch, and composition estimated to be albite (90 percent), quartz (8 percent), tourmaline (2 percent), less

than 1 percent garnet, and a trace of biotite.

Reserves estimated for the entire dike total 30,500 tons of perthite, 7,600 tons of scrap mica, 3,980 tons of soda feldspar, 1,571 tons of lepidolite, 823 tons of topaz, 76 tons of beryl, 282 pounds of monazite, 179 pounds of columbite-tantalite, and 57 pounds of microlite.

Brown Derby No. 5 (pegmatite No. 535)

The Brown Derby No. 5 (pegmatite No. 535) (pl. II) is in a small gulch on the west side of the Brown Derby ridge at an elevation of 8,900 feet. It is in the south-central part of sec. 34, T. 50 N., R. 3 E., New Mexico principal meridian.

This claim is reached by a short spur road from the main Brown Derby road and is owned by Mrs. Marie Disberger. The workings consist of two small open cuts and an adit. The larger cut has a main part, 32 feet long, 15 feet wide, and 18 feet deep at the eastern face. The southern branch of this cut is 12 feet long, 6 feet wide, and 6 feet deep at the northeast face. An adit, approximately 10 feet long, was driven from the eastern end of this cut. A second shallow cut, 10 feet long and 8 feet wide, is northeast of the main cut. This pegmatite was originally mapped by J. B. Hanley and R. Miller III, on September 3, 1943, with plane table and telescopic alidade. The internal structure of the pegmatite was revised by the writer in September 1949 (fig. 23).

The Brown Derby No. 5 pegmatite intrudes greenish-black tonalite.

The pegmatite is irregular in shape, has a length of 210 feet, and a maximum width of 50 feet. It consists of three zones: wall zone, intermediate zone, and core. The wall zone is 2.5 feet thick where it is exposed in the large open cut, has a range in grain size of 0.12 inch to 0.25 inch, and

consists of albite (61 percent), quartz (25 percent), perthite (10 percent), muscovite (4 percent), black to greenish-black tourmaline (<1 percent), lepidolite (trace), garnet (trace), and beryl (<0.1 percent). One exposed area of the wall zone, about 3 feet square, contains 17 beryl crystals ranging in size from 0.5 inch by 0.5 inch to 1 inch by 1.3 inches. This unusually rich area averaged 1.13 percent beryl, but contained only three cobbable crystals. The beryl is in blue-green euhedral crystals.

Inside the wall zone a small intermediate zone is found in the southeastern part of the pegmatite. This zone is approximately 50 feet long and 14 feet thick. It has an average grain size of 4 inches and consists of quartz (55 percent), white cleavelandite (35 percent), lepidolite (5 percent), white blocky perthite (4 percent), muscovite (1 percent), beryl (0.1 percent), topaz (<1 percent), garnet (<1 percent), greenish-black tourmaline (<1 percent), apatite (trace), microlite (trace), and columbitetantalite (trace). Lepidolite occurs as fine-grained aggregates and as large flat sheets that are 4 to 6 inches in diameter. The workings are on this zone, and mining was directed toward the recovery of the fine-grained lepidolite; most of the large sheets were thrown onto the dump. The beryl is in blue-green euhedral crystals from 0.5 to 3.5 inches in diameter. The beryl content increases in the northern part of the zone and appears richest in the small northern pit. A beryl count made on an area 4 feet by 5 feet in this upper pit contained 0.43 percent beryl. About 15 pounds of beryl were also found lying on the dump from this pit. The topaz is milky white and forms euhedral crystals 4 to 6 inches long adjacent to the lepidolite. The garnet forms crystals as much as 1.5 inches in diameter, and commonly is surrounded by coronas of muscovite. It is most common near the contact

with the underlying wall zone. The apatite is in widely scattered lightblue 0.5 inch crystals. Microlite was not seen in place, but was found on the dump in distorted octahedrons 0.12 to 0.25 inch in diameter. The olive-green microlite is faintly radioactive and is found between plates of cleavelandite. One crystal of columbite-tantalite, 0.06 inch thick by 0.75 inch wide, was found between plates of cleavelandite.

The core at its southeastern end has a gradational contact with the intermediate zone. The minerals of the core have an average grain size of 6 inches. It consists of white massive quartz (40 percent), white blocky perthite (39 percent), albite (20 percent), muscovite (1 percent), lepidclite (trace), blue-green beryl (1 crystal, 4 by 6 inches), and columbite-tantalite (two thin pieces, 0.5 inch long by 0.06 inch thick).

The reserves of the intermediate zone are estimated to be 0.75 tons of beryl, 30 tons of perthite, 7.5 tons of scrap mica, 260 tons of cleave—landite, and 37 tons of lepidolite.

Pegmatite No. 537

Pegmatite No. 537 (pl. II) is on a small ridge in the SE¹/₄ sec. 34, T. 50 N., R. 3 E., New Mexico principal meridian. It contains one small adit approximately 4 feet wide and 6 feet long. This working is several hundred feet above and 600 feet to the northeast of the Brown Derby mine road. An old unidentified location marker is found near the adit.

Pegmatite No. 537 is an irregular dumbbell-shaped body (fig. 24) 530 feet long and 84 feet in maximum width. The pegmatite intrudes hornblende gneiss, has a fine-grained wall zone, and has three small discontinuous core segments. The wall zone is more than 90 percent of the pegmatite and has an average grain size of 0.25 inch. It consists of albite (60 percent), quartz (25 percent), perthite (10 percent), muscovite (5 percent), and

garnet (less than 1 percent).

The northern and southern core segments of this pegmatite have an average grain size of 6 inches and consist of perthite (66 percent), quartz (20 percent), albite (10 percent), and muscovite (4 percent). The perthite ranges in size from 3 to 12 inches. Beryl is found in only the northern core segment. A grain count on the sides of the small adit in this core showed 0.31 percent beryl. The beryl is present as 22 pale green euhedral crystals that range in size from 0.25 by 0.12 inch to 2.5 by 4 inches. The core segment containing beryl is 52 feet long and 10 feet wide, and is estimated to contain 3.4 tons of beryl. The central core segment is a few feet wide, has an average grain size of 6 inches, and contains quartz (85 percent), perthite (10 percent), albite (5 percent), and muscovite (less than 1 percent). No beryl was noted in this pod.

Pegmatite No. 538

Pegmatite No. 538 (pl. II) caps the top of a small ridge in the SE¹/₂ sec. 34, T. 50 N., R. 3 E., New Mexico principal meridian. One small cut, 6 feet square by 1 foot deep, exposes the southern end of this pegmatite. This cut is approximately 650 feet northeast of the Brown Derby road and several hundred feet above it. Pegmatite No. 538 is an elongate lenticular-branching pegmatite, approximately 550 feet long and 60 feet in maximum width (fig. 24). The pegmatite intrudes hornblende gneiss, and consists of wall zone and three small discontinuous core segments located in the thicker parts of the pegmatite. The wall zone comprises more than 60 percent of the pegmatite, has an average grain size of 0.25 inch, and consists of albite (57 percent), quartz (25 percent), perthite (10 percent), and muscovite (8 percent). The core segments have an average grain size of 4 inches and consist of quartz (50 percent), perthite (32 percent),

tals 6 to 8 inches in diameter. Pale-green beryl was noted only in the southernmost core segment, where it was estimated, from several beryl counts, to be 0.95 percent of the rock. The beryl crystals range from 0.12 inch by 0.12 inch to 6 inches by 6 inches in area. The southern pod is 128 feet long and has a maximum width of 35 feet. It was calculated to contain 42 tons of beryl to a depth of 24 feet.

Pegmatite No. 560

Pegmatite No. 560 (pl. II) is an unclaimed pegmatite at the foot of the mountains on the east side of Quartz Creek in the west central part of sec. 34, T. 50 N., R. 3 E., New Mexico principal meridian. This pegmatite is 1,500 feet south of State Highway 162 and lies directly across a meadow. It is extremely irregular (fig. 25) and has a maximum length of 430 feet and a maximum width of 200 feet. It cuts across the hornblende gneissgranite contact. This pegmatite consists of a narrow wall zone, a large core, and a small pod on the south end. The wall zone has an average grain size of 0.12 inch and is estimated to consist of albite (60 percent), quartz (36 percent), perthite (4 percent), muscovite (less than 1 percent), and garnet (trace). The core comprises the greater part of the pegmatite and has an average grain size of 4 inches. It consists of perthite (50 percent), quartz (30 percent), albite (20 percent), and muscovite (trace). On the south end of the pegmatite there is a lenticular pod. 78 feet long and 18 feet wide. This pod has an average grain size of 1 to 2 feet, and contains perthite (75 percent), quartz (20 percent), albite (5 percent), and beryl (0.45 percent). The beryl is pale green and ranges in size from 1 by 2 inches to 4 by 8 inches. This pod contains the only beryl that was

noted in the pegmatite. Reserves of beryl are calculated to be 9.5 tons to a depth of 20 feet.

Beryl and Mare Minerals Lode (pegmatite No. 590)

The Beryl and Rare Minerals Lode (No. 590, pl. II) is a small lenticular pegmatite on the north-facing slope of Tollgate Gulch, in the SE₄ sec. 34, T. 50 N., R. 3 E., New Mexico principal meridian. This lenticular pegmatite is 154 feet long and 55 feet wide, and dips gently to the south at the angle of 5 to 10 degrees (fig. 25).

The property is about a quarter of a mile south of a small private road in the bottom of Tollgate Gulch, and is reached by a narrow path winding up the hillside.

The claim on this pegmatite was located by Jesse Field on November 27, 1949. Mr. Field has opened at least six small pits ranging from a few feet square by 1 foot deep, to 22 feet long, 10 feet wide, and 2 feet deep. These pits are on local concentrations of beryl and thus expose the richest parts of the beryl-bearing pegmatite. To date Mr. Field has recovered approximately 480 pounds of beryl, 2 pounds of columbite-tantalite, and approximately 800 pounds of muscovite. The Beryl and Rare Minerals pegmatite intrudes hornblende gneiss and may be divided into three zones: wall zone, intermediate zone, and core. The top of the pegmatite has been eroded, exposing the flatlying central units. The wall zone is thin and irregular, and is exposed in only a few places along the edge of the pegmatite. It has an average grain size of 0.25 inch and consists of albite (55 percent), perthite (20 percent), quartz (25 percent), and muscovite (< 1 percent).

The intermediate zone is well exposed by the workings. It has an average grain size of approximately 3 feet and is estimated to consist of

perthite (50 percent), muscovite (30 percent), quartz (20 percent), albite (<1 percent), beryl (0.1 percent), columbite-tantalite (<0.05 percent), gahnite (traces), and an unidentified mineral which resembles the samarskite, fergusonite, or euxenite group of minerals. The perthite occurs in crystals from 1 to 5 feet in diameter. The muscovite is abundant in the outer part of this zone and occurs in books as much as 8 inches across. It is reeved, soft, and heavily stained and is all scrap mica. It closely resembles the mica obtained at the Bucky and Buckhorn properties. Beryl ranges from 0.5 to 8 inches in diameter and is white. The percentage of beryl obtained in the pits is estimated from the amount of beryl recovered and the size of the workings to be 0.4 percent. Because the pits were in the beryl-rich parts of the pegmatite and as many parts of this zone are completely barren of beryl, the overall content in this zone is approximately 0.1 percent. Columbite-tantalite was found intergrown with perthite in one pit. These crystals are from 0.01 to 0.12 inch thick and as much as 2 inches across, but no columbite-tantalite is exposed in the rest of the pegmatite. Gahnite is intergrown with fine muscovite in one small area. This mineral crystallizes as dark-green octahedrons 0.01 inch in diameter. Inside the intermediate zone is a core made up entirely of quartz, that extends the length of the pegmatite.

White Spar No. 2 (pegmatite No. 604)

The White Spar No. 2 pegmatite is in sec. 35, T. 50 N., R. 3 E., New Mexico principal meridian. It is on the north side of Tollgate Gulch, 0.9 mile from State Highway 162, and is reached by a mine road that follows the gulch. The pegmatite is now being mined for lepidolite by Compsolidated Feldspar Company. It was located in August 1942 and is owned by the Colorado Feldspar Company.

The pegmatite was examined and mapped with plane table and telescopic alidade by E. W. Heinrich and Roswell Miller III, on July 28, 1943 /.

_____ Hanley, J. B., Heinrich, E. W., and Page, L. R., Pegmatite investigations in Colorado, Wyoming, and Utah, 1942-1944: U. S. Geol. Survey Prof. Paper 227, pp. 77-80, 1950.

Two prospect pits have been made in the pegmatite, one about 40 feet long and 10 feet wide at the north end of the dike and one approximately 60 feet long and 25 feet wide at the south end.

The pegmatite is about 260 feet long and ranges in width from 6 feet near the center to nearly 50 feet at the north end. The trend is north, but the southern contact of the pegmatite strikes N. 150 W., and dips 700 NE. The pegmatite cuts hornblende gneiss, whose foliation strikes N. 250 W. and dips 70° to 80° NE. The pegmatite consists of a core of fine-grained albite-quartz-perthite-lepidolite pegmatite surrounded by a discontinuous wall zone of albite-perthite-quartz-muscovite pegmatite. The wall zone is discontinuous over the length of the pegmatite, and has a maximum width of 5 feet. Its grain size averages 0.25 inch, and the composition is estimated to be 45 percent albite, 30 percent perthite, 20 percent quarts, 5 percent moscovite, and less than I percent lepidolite. The lepidolite has an average grain size of 0.12 inch but occurs in books as much as 1 inch across. The core ranges from 6 to 32 feet wide. It has a finer grain size than the wall zone, averaging 0.02 inch, and contains 45 percent plagioclase, 35 percent quartz, 10 percent perthite, 10 percent lepidolite, less than I percent garnet, and traces of beryl, microlite, fluorite, and chrysocolla. The lepidolite has an average grain size of 0.03 inch, and occurs in lenses and stringers up to 4 inches wide. The lepidolite exposed in the northern pit is banded with coarser-grained albite and quartz.

A grab sample of the core taken by Heinrich ___ from the southern pit

and analyzed spectrographically by the Geological Survey contained 0.7 percent Li₂0, or about 17 percent lepidolite, 0.05 percent BeO, and no Cb or Ta.

White Spar No. 1 (pegmatite No. 636)

The White Spar No. 1 pegmatite is in sec. 35, T. 50 N., R. 3 E., New Mexico principal meridian. It is on the north side of Tollgate Gulch and is connected to State Highway 162 by a mine road 0.7 mile long running up the bottom of the gulch. A claim was located on this pegmatite in 1942 by the Colorado Feldspar Company. E. W. Heinrich and Roswell Miller III of the Survey examined and mapped this property with plane table and telescopic alidade on July 29, 1942.

_/ Hanley, J. B., Heinrich, E. W., and Page, L. R., Pegmatite investigations in Colorado, Wyoming, and Utah, 1942-1944: U. S. Geol. Survey Prof. Paper 227, pp. 77-80, 1950.

The mine workings consist of five prospect pits, the largest of which is 50 feet long and has a maximum width of 25 feet.

The pegmatite crops out on the top and on the south-facing slope of a narrow ridge paralleling and separating Tollgate Gulch from a gulch to the north. The pegmatite is intruded into hornblende gneiss, but none of the contacts are exposed.

The pegmatite trends N. 20° E., and dips 30°-35° SE. It has a length of 200 feet and a maximum width of 85 feet.

Four zones are well developed within the pegmatite. A wall zone of fine-grained albite-perthite-quartz-muscovite pegmatite completely surrounds an intermediate zone of fine-grained cleavelandite-quartz-perthite-lepidolite pegmatite and cores of lepidolite-quartz pegmatite and quartz pegmatite. The wall zone is 9 feet thick on the hanging-wall side and 33 feet on the footwall side $\sqrt{\ }$, and has an average grain size of 2 inches.

J Hanley, J. B., Heinrich, E. W., and Page, L. R., op. cit., p. 77,

The composition is estimated to be 45 percent albite, 32 percent perthite, 20 percent quartz, and 3 percent muscovite. Perthite is in crystals as much as 24 inches long and 15 inches wide, and muscovite books average 1.5 inches across and 1 inch thick.

The cleavelandite-quartz-perthite-lepidolite intermediate zone on the western and southern edges of the quartz core, is 90 feet long and ranges in width from 1 to 18 feet. The grain size is 1 inch, and the zone is estimated to contain 55 percent cleavelandite, 25 percent quartz, 15 percent perthite, 5 percent lepidolite, 0.01 percent beryl, 0.003 percent topaz, and 0.0003 percent columbite-tantalite, and a trace of microlite. Perthite crystals average 12 inches in length and 8 inches in width. The lepidolite is in books 3 inches across and 0.5 inch thick. The beryl is yellow to pale blue-green and occurs in crystals from 0.5 to 1.75 inches in diameter. Topaz crystals are small, ranging in size from 0.06 to 1 inch. The columbite-tantalite crystals are as much as 0.4 inch long and 0.25 inch wide.

The pegmatite cores are of two types: white massive quartz pegmatite and lepidolite-quartz-microlite pegmatite. The quartz pegmatite occurs as one large irregular mass 80 feet long and 3 to 13 feet wide, and seven

smaller lenses.

The core of the lepidolite-quartz-microlite pegmatite is on the west side of the quartz pegmatite, between the core and the intermediate zone of cleavelandite-quartz-perthite-lepidolite pegmatite. This core is 20 feet long and from 1 to 8 feet wide, and has a grain size of 0.03 inch, and contains approximately 90 percent lepidolite, 10 percent quartz, and 0.1 percent microlite.

Buckhorn (pegmatite No. 659)

The Buckhorn (pegmatite No. 659, pl. II) caps the top of a ridge on the north side of Tollgate Gulch in the SE¹/₄ sec. 27, T. 50 N., R. 3 E., New Mexico principal meridian. This irregular pegmatite (fig. 26) has a maximum length of 1,750 feet and a maximum width of 1,360 feet. It is between 8,900 and 9,400 feet above sea level and 350 to 850 feet above Tollgate Gulch. The nearest road is State Highway 162, 0.4 mile to the west.

At least three claims have been located on this pegmatite. Claim notices show two of these to be: the Buckhorn, which is on the northwestern part of the pegmatite and the Feldspar claim in the northeastern part of the pegmatite, both located by Bert and Florence Tucker. On the Buckhorn claim there are several small trenches, the largest is 30 feet long and 5 feet wide, and the smallest is about 4 feet long, 3 feet wide, and 2 feet deep. The Feldspar claim has a trench 15 feet long, 3 feet wide, and 2 feet deep. Several hundred feet to the east of the Feldspar claim is an unnamed claim which has a small shaft, 4 feet square and 8 feet deep, and a trench 15 feet long and 3 feet wide. No mining has been done on these claims.

The Buckhorn pegmatite (fig. 26) intrudes hornblende gneiss and tonalite. The greater part of this pegmatite has only one zone, but it contains several small disconnected core segments in its upper part along the ridge. Around one of these cores is a small intermediate zone. The small cuts in the Buckhorn claim are made on this intermediate zone. The wall zone, which forms more than 90 percent of the pegmatite, has an average grain size of 0.25 to 0.5 inch. It is estimated to contain albite (59 percent), quartz (20 percent), white to pink perthite (20 percent), fine-grained graygreen muscovite (1 percent), garnet (trace), and biotite (trace). Though most of these minerals are fine grained, the perthite occurs in blocky crystals 1 to 3 inches in diameter.

The intermediate zone is 1.5 to 2 feet thick. It has an average grain size of approximately 1 foot, and consists of perthite (50 percent), muscovite (40 percent), and quartz (10 percent). A few greenish crystals of beryl, approximately 1 inch in diameter, were noted. Adjacent to the core in a feldspar-rich part of this zone three or four crystals of columbite-tantalite and monazite, about 0.5 inch long, were found. The muscovite in the intermediate zone occurs in books as large as 10 by 18 inches. It is greenish-gray, and has a strong "A" structure. It crumbles easily between the fingers. It closely resembles the scrap mica from the Bucky mine, which is prized as a grinding mica. The intermediate zone is about 150 feet long, and diminishes in grade to the south. The amount of scrap mica available, therefore, is small.

The core segment inside the intermediate zone is south of the other core segments and differs considerably from them in composition and texture. This core has an average grain size of 2 feet and consists of perthite (91 percent), quartz (7 percent), muscovite (1 percent), and

beryl (0.7 percent) as determined by a beryl count along the sides of the largest cut. The beryl is pale green and ranges in size from 0.017 by 0.058 foot to 0.27 by 0.45 foot.

The other cores lying along the top of an east-west trending ridge are only 10 to 20 feet thick. They may be the eroded remnants of a once much larger and continuous core. The average grain size of the minerals in these core segments is 8 to 12 inches, and though they vary in the percentage of minerals, they are estimated to contain perthite (53 percent), quartz (45 percent), albite (1 percent), and muscovite (1 percent).

The small shaft to the east of the Feldspar claim was sunk in a quartz-rich part of this pegmatite. It yielded approximately 75 pounds of beryl. These beryl crystals are 1 to 2 inches across, are white, and closely resemble quartz. This appears to be a beryl-rich pocket and others might be found on further exploration. The amount of beryl is not likely to be large, as the core segments are extremely thin. A limited amount of feldspar could be obtained from these core segments, but as the cost of feldspar is low and transportation costs are high, the economic feasibility of this is questionable.

Black Wonder (pegmatite No. 847)

The Black Wonder pegmatite is the largest pegmatite in the district, covering parts of secs. 20, 21, 22, 27, 28, 29, 32, and 33, T. 50 N., R. 3 E., New Mexico principal meridian. It is 12,600 feet long and has a maximum width of 6,700 feet. The northeast end is less than 200 feet from Willow Creek and the southwest end is at Big Gulch. A road from Big Gulch to the State Highway 162 traverses the pegmatite for 1.5 miles. Much of the eastern edge is within a quarter of a mile of the highway. The western part

of the pegmatite forms the southern extension of a prominent north-trending ridge. This ridge rises to the north, and the highest point on the pegmatite is over 9,700 feet. The southern and eastern edges of the pegmatite are at an altitude of slightly above 8,500 feet.

The Black Wonder is extremely irregular, consisting of a large number of intersecting dikes of uneven spacing and size.

Most of the pegmatite intrudes hornblende gneiss, but part of it cuts coarse-grained granite at the southeast and pre-Cambrian sediments in a small area at the northeast.

The pegmatite has aroused little mining interest. Two claims have been filed on different parts of the pegmatite: the Black Wonder and the Beryl claim. The Black Wonder prospect, in the eastern part of sec. 29, was located in May 1948 on a magnetite-rich area in the pegmatite, and consists of one small pit. The Beryl claim, located in June 1948 by Bert Tucker, is in sec. 27 and consists of three pits on a beryl-bearing unit and a fourth on a monazite-bearing unit.

The pegmatite consists of a thick wall zone enclosing small widely scattered cores and cut by occasional fracture fillings. Only a few of the cores have an intermediate zone between them and the wall zone. The wall zone, constituting over 95 percent of the pegmatite, varies in texture and composition. In the southern and western parts of the pegmatite it is a graphic granite unit, grading to the north and northeast into a unit with only a few crystals of graphic granite in a matrix of albite and quartz. The wall zone at the south and west end has an average grain size of 3 inches and is estimated to contain 60 percent perthite, 24 percent albite, 15 percent quartz, 1 percent martite, less than 1 percent biotite, and a trace of garnet. Graphic granite crystals, as much as 5 feet across,

constitute 50 percent of this part of the wall zone. Small local concentrations of martite are common, such as the one upon which the Black Wonder claim was made. Martite comprises 10 percent of the wall zone at the prospect pit and to the northeast for 50 feet, in crystals as much as 4 inches across. The wall zone to the north and east has an average grain size of 1.5 inches and is estimated to contain 55 percent albite, 30 percent perthite, 15 percent quartz, less than 1 percent garnet, and a trace of martite. Graphic granite constitutes less than 10 percent of this part of the wall zone in crystals less than 6 inches across.

At the Beryl claim, two different types of intermediate zone surround quartz cores. One intermediate zone contains monazite and columbitetantalite, and the other beryl crystals. The intermediate zone that contains beryl is between two small quartz pegmatite cores, 7 feet apart, the larger of which is 25 feet long by 10 feet wide. Two prospect pits, the larger of which is 9 feet long by 6 feet wide, are along the east side of the larger quartz pegmatite core. Another pit, 15 feet long and 4 feet wide, is on the east edge of the smaller quartz pegmatite core. The intermediate zone is not exposed completely around the larger core, but lies east of it, surrounding the smaller one. This zone has a maximum size of 30 feet long and 15 feet wide. The grain size averages 2 inches, and the zone is estimated to contain 79 percent albite; 10 percent quartz, 5 percent perthite, 5 percent muscovite, 1 percent garnet, and 0.2 percent beryl. The muscovite is in books from 0.25 to 5 inches across. Garnet crystals range in size from 0.25 to 2 inches. Beryl is concentrated along the eastern edge of the larger quartz pegmatite core in semitransparent yellowishgreen crystals from 0.5 to 1.25 inches in diameter.

The monazite-bearing intermediate zone is approximately 400 feet north-

east from the beryl-bearing zone, and is exposed by a prospect pit 4 feet long and 3 feet wide, on the east side of a quartz pegmatite core. The core is 15 feet long, 6 feet wide, and is 2 feet thick at its edge in the pit. This intermediate zone is estimated to be 15 feet long and 4 feet wide. Its grain size is 6 inches and the composition is estimated to be 55 percent albite, 30 percent muscovite, 15 percent quartz, and traces of monazite and columbite-tantalite. Muscovite books are as much as 8 inches across, and monazite crystals average 0.75 inch long by 0.12 inch wide. The columbite-tantalite averages 0.12 inch by 0.06 inch.

The cores are nearly all small, measured in tens of feet long and less than 10 feet wide. They range from 100 percent quartz to 10 percent quartz and 90 percent perthite. In the northeast part of the pegmatite muscovite books form as much as 30 percent of the cores in crystals as much as 5 inches across. In some of the cores the quartz is smoky, suggesting the presence of radioactive minerals. Several small crystals of allanite were found in one core.

In various places the wall zone is cut by fracture fillings of white quartz. These fracture fillings range from a fraction of an inch to 6 inches in width.

Trio No. 1 (pegmatite No. 1402)

The Trio No. 1 pegmatite is on the ridge west of Willow Creek at an altitude of 10,000 feet, in secs. 16 and 21, T. 50 N., R. 3 E., New Mexico principal meridian. The nearest road is along Willow Creek 1 mile northeast of the claim. This road joins State Highway 162, 2.5 miles to the southeast. The claim was located on May 2, 1949 by Bert Tucker, George Tucker, and A. T. Pearson. Discovery workings consist of four small prospect pits,

the largest of which is 13 feet long, 10 feet wide, and 4 feet deep.

The pegmatite is 644 feet long and has a maximum width of 152 feet. It is irregular in shape and is intruded into quartz monzonite. The pegmatite is made up of four zones: a large thick wall zone constituting over 90 percent of the pegmatite, two small intermediate zones, and several small discontinuous cores. The wall zone has an average grain size of 0.75 inch and is made up of albite (45 percent), perthite (40 percent), quartz (15 percent), biotite (less than 1 percent), and martite (less than 1 percent). The intermediate zones are of two types: a quartzalbite-perthite pegmatite and a quartz-albite-muscovite pegmatite, both found around one core. The quartz-albite-perthite pegmatite intermediate zone, 20 feet long by 15 feet wide, is east of the quartz-albite-muscovite pegmatite intermediate zone and separates it from the wall zone. This quartz-albite-perthite pegmatite has an average grain size of 6 inches and is estimated to contain quartz (35 percent), albite (34 percent), perthite (30 percent), garnet (0.5 percent), and biotite (0.5 percent). Six beryl crystals from this zone were found on the stockpile, ranging from 1 to 8 inches in diameter and from 1 to 6 inches in length. The quartz-albitemuscovite pegmatite is 3.5 feet thick and surrounds the core of quartz pegmatite. This zone has an average grain size of 1 inch and is made up of quartz (60 percent), albite (25 percent), muscovite (15 percent), and garnet (less than 1 percent). Muscovite crystals average 1 inch across and 1.5 inches thick.

Cores of quartz pegmatite occur in several places and are as much as 15 feet long by 12 feet wide. The core on which the workings are located is 20 feet long by 5 feet wide and consists entirely of quartz.

Bucky (pegmatite No. 1574)

Introduction

The Bucky pegmatite is an irregular pegmatite on the ridge between Willow and Illinois Creeks. Numerous claims are located on this pegmatite in the E2 sec. 22, T. 50 N., R. 3 E., New Mexico principal meridian. The Bucky claim on which the main workings are found is on its northern end, covering a quartz pod 100 feet long and 60 feet wide. This claim was originally owned by Rod Fields, who has driven several small adits along the southern side of the pod and has produced approximately 17 tons of beryl, 100 pounds of columbite-tantalite, 25 pounds of an unidentified samarskite-like mineral, and 15 pounds of monazite. Scrap mica was at first discarded but approximately 20 tons were stockpiled in September 1948. In the fall of 1948, Mr. Fields sold the property to the Beryllium Mining Company, Inc., which has operated from open pits excavated by blasting and bulldozing. A road was constructed to the mine workings approximately 400 feet above the valley bottom by the Beryllium Mining Company, In May 1950 a small mill for separating the scrap mica was built alongside the mine road. Prior to May 15, 1950, the Beryllium Mining Company, Inc. had produced 32 tons of beryl, 139.6 tons of scrap mica, 1,020 pounds of columbite-tantalite, 15 pounds of monazite, and 13 pounds of a samarskite-like mineral.

The beryl was sold to various buyers in Colorado, and was trucked to Longmont or sold on the property. The scrap mica is sold to Western Non-metallics in Pueblo, Colo., at \$25.00 a ton, delivered in Pueblo. No columbite-tantalite has been sold, and the monazite and the samarskite-like mineral have been sold to Ward's Natural Science Establishment for resale as mineral specimens.

The Bucky mine workings were mapped by Staatz and Flawn between

September 28 and 31 in 1948, with plane table and telescopic alidade.

This map (figs. 27 and 28) covered an area from the northern contact of the pegmatite with the schist to a point 180 feet south of the main quartz pod. A beryl count was made in the mine workings. On November 22, 1949, the map was revised by Staatz and Trites to show the new workings. The outline of the whole pegmatite was mapped by Staatz (pl. II) in the course of regional mapping in September 1949,

Geology

The Bucky pegmatite has been intruded chiefly into hornblende gneiss, but it also cuts several small bands of quartzite. The pegmatite is extremely irregular and contains many small inclusions or pendants of country rock. The main bulk of the pegmatite is a fine-grained discontinuous wall zone and a coarse-grained graphic granite intermediate zone. Inside this are scattered 36 cores of quartz pegmatite segments at least 10 feet long. Some of these core segments are surrounded by one to three intermediate zones. The core segments have a peripheral arrangement in the pegmatite (fig. 27) and some may be fracture fillings rather than true core segments.

The wall zone is discontinuous; it is absent in some parts and several hundred feet thick in others. It has an average grain size of 0.25 inch, and consists of albite (60 percent), perthite (20 percent), quartz (16 percent), muscovite (4 percent), and a trace of garnet.

Inside the wall zone is a thick intermediate zone made up chiefly of graphic granite. The graphic granite aggregates range from 2 inches to 1 foot in diameter, and average about 5 inches. Besides graphic granite, this pegmatite unit contains 3 to 4 percent of cream-colored blocky

perthite, 1 percent of white quartz crystals, 3 percent of fine-grained cream-colored albite, and less than I percent of biotite. The biotite occurs in thin, 6-inch blades that are localized in small areas in this rock. The albite is difficult to distinguish from the perthite, but is most abundant along the contact of the quartz-albite perthite pegmatite. The albite has a minimum index of refraction (N_{47}) of 1.530 \bullet 0.002. The estimated bulk composition of this rock is perthite (77 percent), quartz (20 percent), albite (3 percent), biotite (1 percent). Quartz albite pegmatite is the most common type of pegmatite adjacent to and encircling quartz pods. Some quartz pods have no other intermediate zones separating them from the graphic granite pegmatite, while others have as many as three. Good exposures of this zone are found around the large pod on the Bucky claim. The quartz-albite pegmatite has an average grain size of 0.5 inch, and usually contains equidimensional quartz grains surrounded by albite. The estimated composition of this rock is quartz (55 percent), albite (40 percent), perthite (3 percent), muscovite (1 percent), and garnet (1 percent). The albite is cream-colored and usually fills around the quartz crystals. It has a minimum index of refraction (Na1) of 1.532 * 0.002. The perthite is commonest near the perthite-quartz pegmatite zone where it occurs as graphic granite, and near the muscovitefeldspar-quartz-beryl zone where it occurs as cream-colored blocky crystals about 4 inches across. The muscovite is in light-colored irregular books 0.25 to 0.75 inch long. It occurs in local aggregates, comprising as much as 10 percent of the rock. Adjacent to the core of the Bucky mine the feldspar is considerably kaolinized.

Muscovite-feldspar-quartz-beryl pegmatite predominates around the large Bucky core segment (figs. 27 and 28) but is also well developed

around at least two other core segments and may be present to a minor extent around several more. This zone weathers easily, and is usually concealed by quartz float. It has been well exposed along the southern and eastern side of the Bucky claim, where mine faces are over 20 feet high, This zone extends around three-fourths of the Bucky core segment but pinches out in the northwest quarter and is from 1 to 10 feet thick. muscovite-feldspar-quartz-beryl pegmatite zone has a grain size which ranges from 3 inches to over 8 feet, and has an average of about 2 feet. It has an estimated composition of muscovite (40 percent), feldspar (31 percent), quartz (20 percent), beryl (8.9 percent), columbite-tantalite (0.11 percent), samarskite-like unidentified mineral (0.003 percent), and monazite (0.003 percent), topaz (< 1 percent), gahnite (< 1 percent), phosphates (trace), and lepidolite (trace). Muscovite makes up from 10 to 80 percent of the rock and is found in books as much as 1 foot across; the average is 6 inches. The books are heavily lined, have irregular surfaces, contain minute crooked fractures, and have a prominent "A" structure. Both red and black mineral staining is common. This mica is also scrap and is quite soft, which makes it an excellent grinding mica. The feldspar occurs chiefly as cream-colored blocky perthite and as cream-colored finegrained albite. The albite is commonest in heavy muscovite concentrations and has a minimum index of refraction (Na) of 1.531 ± 0.002. Because of the heavy kaolinization of both feldspars, the relative proportions of perthite to albite could not be readily determined. Quartz occurs as large white crystals several feet in diameter. Beryl is found in large white to pale-green euhedral crystals. A total of 64 beryl crystals was noted in 344.5 square feet of muscovite-feldspar-quartz-beryl pegmatite measured along the mine walls. The crystals ranged in area across their

bases from 0.007 to 5.0 square feet, and averaged 0.70 square feet. Beryl is more common and occurs in larger crystals in the perthite-quartz-rich part than in the muscovite-rich part. A beryl count made in Fields' early workings, which were driven on a beryl-rich concentration, gave an average of 13 percent beryl. Since that time most of this zone has contained much less beryl. A second pocket was opened in April 1950, from which approximately 9 tons of beryl had been taken prior to June of the same year. The beryl in this, as in all zones, is concentrated in pockets separated by almost barren rocks. The beryl has a maximum index of refraction (N_{o}) of 1.578 ± 0.002, which corresponds to approximately 13.2 percent BeO. A small part of the beryl has been kaolinized. The columbite-tantalite, monazite, and an unidentified samarskite-like mineral usually occur together in erratic pockets. They were found in some of the early workings adjacent to the core segment. The columbite-tantalite occurs in black tabular crystals as much as 6 inches across. Monazite occurs in 0.25 to 1 inch long reddish-brown euhedral crystals and the adjacent feldspar is frequently stained red. The samarskite-like unidentified mineral has been found in masses as much as 5 inches across. It is dark greenish-black in color, has conchoidal fracture, a greasy luster, and no apparent crystal form. This mineral is metamict and its X-ray pattern does not agree with that of samarshite, fergusonite, euxenite, allanite, or uraninite. A more complete discussion of this mineral is given in the section on mineralogy.

Topaz has been reported from the Bucky core segment, but is found in greater abundance around a small pod on the southwestern part of the pegmatite. The zone in this area may contain as much as 1 percent of topaz in crystals 1 to 4 inches across. Lepidolite has been found adjacent to the topaz but is quite rare and in very fine-grained aggregates.

Lithiophilite-triphylite has been found in a few rare crystals adjacent to the Bucky core segment and another small pod in the extreme northeastern end of the pegmatite.

In addition to the muscovite-feldspar-beryl pegmatite, there is a coarse-grained perthite pegmatite around the Bucky core segment. This rock has an average grain size of 6 feet and is estimated to consist of 93 percent of cream-colored blocky perthite, 7 percent of quartz, and less than 1 percent of albite and muscovite. The perthite is slightly kaolinized. About 30 tons of cream-colored perthite has been stockpiled at the mine, but none had been sold prior to June 1950.

Quartz core segments are found scattered throughout the pegmatite and range from a few feet long and less than 1 foot wide to the one 100 feet long and 80 feet wide on which the Bucky claim is located. The quartz pegmatite is made up of 100 percent white massive quartz. As this rock is resistant to erosion, it forms prominent knobs, and joint blocks commonly cover the adjoining pegmatite.

A unit which is believed to be a fracture filling was found in two places along the outer edge of the pegmatite. This unit has an average grain size of 3 feet and is estimated to consist of 50 percent quartz and 50 percent perthite,

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humber and name of peg-	Type and	Alteration	Relation to	Shape	Internal	Texture							Service.		0.0		Mi	neralo	еу							
natite (Pl.2)	formation		wall rock		structure	(inches	Plagi	oclase	Per	thite	Grap		Que	artz	Musc	ovite	Gar	net	Tours	aline	Lithium	minera	als.	Other	mineral	
							Per-	Size	Per-	Size	Per-	Size inches	Per-	Size inches	Per-	Size inches	Per-	Si ze inche	Per-	Size	Mineral	Per-	Size	Mineral	Per- cent	(inche
1	Tenalite.		Not expessed.	Lenticular	One unit .	3/4	45		35				25		< 1											
2	De.		Do.	Do.	Do.	2-3	15		49				35		1											
3	De.		Do.	Oval.	Do.	1/4-1/2	47		25				25		3		< 1									
4	Do.		De.	Lenticular	. Do.	3/4-1	53	1/16	15				30	12	1		1	1/8						Biotite. Gahnite.	Trace.	1/32
5	Tenalite and coarse-grained	Hone .	Cross-sutting	Lenticular	De.	1/4	54		20				25		1		< 1							Beryl.	4Crystal	. 1/8-
	granite.			branchis	E. Do.											4,										
6	Granite.	De.	1	Lenticular	. Do.	1/8-1/4	65		15				20		< 1	•	race.									
7	fonalite and granite.		Not exposed.	Lenticular branching.	Do.	1/8-1/4	58		20				20		2		< 1									
8	Bietite gneiss and granite.	None.	Conformable.	Lenticular branching.	Do.	1/4	60		20				20		< 1		Trace									
9	Tonalite and granite.	De.	Oress-sutting	Lenticular	. Do.	1/4-1/2	62		18				20		< 1		< 1									
10	Unknown.			Do.	Do.	1/16	69		10				20		Prace.		< 1									
n	Tonalite.	Jone.	Cross-outting	Lenticular branching.	Do.	1/4	60		20				20		< 1		< 1	.5								
	Tonalite, coarse- grained granite, fine-grained granite.	Do.	Do.	Do.	Do.	1/4-1/2	14		35				20		< 1		< 1									
13	Coarse-grained granite.			Leaticular	. Do.	1/8-1/4	65		15				20		< 1		Trace									
14	Tonalite.		Cress-outting	Leaticular	. Do.	1/8-1/4	65	1	15				20		< 1		< 1		Trace							
15	Pine-grained granite.			Do.	Do.	1	30		35				35		2		< 1									
16	Tonalite.		Bot expessed.	Lenticular	De.	1/2-3/4	50		15				35		< 1		41									
17	Unknown .		De.	Leaticular branching.	De.	1/2-3/4	55	1/8	25	4-6			20	4-6	<1		L1									
16	Tonalite and coarse-grained granite.	Hone.	Cm so-sutting	. Do.	Do.	3-4	15	1/8	59	8-10			25	5-10	1	3-4	< 1									
19	Tonalite.		Bet exposed.	Lenticular	. Do.	4	25		50	-		100	25				Trace									
20	Do.	Fone.	Cress-outting	.Irregular	. Do.	1/8	72		7	1873			20		< 1	12	< 1									
n	Do.		Bot exposed.	Lenticular	. Do.	1/2	55		25			7 . 10	20		< 1		< 1							Columbite	1 Crysta	

	Wall r	ock													Pog	matite			_							
Sumber and name of peg-	Type and	Alteration	Relation to	Shape	Internal structure	Texture											M	neralo	EY							
(Pl. 2)	formation		wall rock		structure	1EGES)	Plagio	clase	Perti	hite	Grap		Qua	rts	Musc	ovite	Gar	net	Tours	aline	Li thium	miners	g.	Other	mineral	10
								Sise												Sise	Mineral	Per-	Size	Mineral	Per- cent	Sise (inches
k 2	Hernblende gneise.		Bot exposed	Lentiqual x branchiz		1/8	71						20		1		Trace.									
42	Do.		Do.	Irregular.	Do.	1/16-1/8	74		10				15		1		1									
43	Coarse-grained granite.	Sone.		Do.	Do.	1/8-1/4	46		30				80		2		1									
14	Do.			Lenticular	. Do.	1/8	74		5				20		1		1									
45	Hornblende gneiss.		Not exposed.	Do.	Do.	1/4	65		15				19		1											
46	Do.		Grossoutting	.Lenticula branchin		1/8	74		5				20		1		1									
47	Hernblende gneiss and granite.		Do.	Do.	Do.	1/8-1/4	65		15				20		<1		Trace.									
46	Hornblende gneiss.		Do.	Do.	Do.	1/8-1/4	69		10				20		1		<1									
49	Do.		Do.	Do.	Do.	1/16-1/0	64		10				25		1		<1							1		
50	Do.		Not exposed	Lenticula	r. Do.	1/4	65		15				20		<1	-	Trace.	/								
51	Do.		Do.	Irregular.	Do.	1/4	65		15				20		Trace		Trace.									1.00
52	Tonalite		Grossoutting.	Lenticular branching		1/8-1/4	70		7				23		Trace		3/2									
53	Hornblende gneiss.			Lenticular	Do.	1/32	80		5				15		Trace		<1									
54	Do.		Not exposed.	Do.	Do.	1/4	55		20				25		< 1		<1									
55	Tonalite		Orossentting.	Irregular.	Borth end.	1/4	59 27		27				20 45		<1		Trace.							Beryl.	1 Crystal	1-1/8
56	Do.		Not exposed.	Lenticular	One unit.	1/8	75		5				50		< 1		<1									
57	Hornblende gneiss.		Do.	Do.	Wall some Cores	1/8-1/4	70 25		15 35				15		Trace		Trace.									
58	Tenalite		Do.	Do.	One unit.	1/16-1/	77		3				20		Trace		Trace.	187								
59	Do.		Do.	Do.	Do.	1/8	80		5				15				Trace.			- 3			200			
60	Do.		Do.	Irregular.	Do.	1/8	71		8		3/44		20		1		1									
61	Do.	CAN'S	Do.	Leaticular	Do.	1-2	35		30		line.	12.11	35		1											1
62	Do.		Do.	Do.	Wall some Core	1/8	70 5		15 10				15 85		1		1	9								
63	Do.		Crossoutting	Lenticular branching		1	54		20				25		1		Trace.									

	Vall r	ock													Pog	matite						-				
nber and ne of peg-	Type and	Alteration	Relation to	Shape		Texture											M	neralo	ey .							
tite Pl. 2)	formation		wall rock		structure		Plagio	clase	Pertl	hite	Grap		Qua	rts	Musc	covite	Gar	net	Tours	aline	Li thium	miner	als	Other	minerals	
							Per-	Size Inches	Per-	Size nches	Per-	Size inches	Per-	Size Inches	Per-	Size	Per-	Size inches	Per-	Size	Mineral	Per-	Size	Mineral	Per-	Si (inc)
64	Tonalite.	Hone.	Crosscutting.	Lenticular branching	-Wall some West branc Core West branc Hast branc	6	69 15		15 43 25				15 40 20		1 2 <1		<1							Columbite- tantalite Honasite Fourmeline.	Grystal	
65	Do.		Not exposed.	Do.	One unit	1/8	69		10				20		1		<1									
66	Do.		Do.	Lenticular	Do.	1/2-1	45		10				45		<1		Trace.									
67	Coarse-grained granite.		Do.	Do.	Do.	1/4	65		12				23		<1		<1									
68	Tenalite.	Service :	Do.	Do.	Do.	1/2-1	50		25				25		<1		Trace.							Beryl.	crystal	8
69	Do.		Crosscutting.	Do.	Do.	1/8-1/4	64		15				20		1		<1									
70	Do.		Do.	Lenticular branching.	Do.	1/8	72		8				20		<1		frace.									
n	Tenalite and coarse-grained granite.	Hone.	Do.	Lenticular.	Do.	1/4-1/2	50		25				25		<1		Trace.									
72	De.	Hone.	Do.	Lenticular- branching.	Do.	1/5-1/4	65		15				50		<1		<1									
73	Tonalite		Set exposed.	Lenticular	Do.	1/8-1/4	55		20				25		<1		Prace.									
74	Do.		Crosscutting.	Lenticular- branching.	Do.	1/8-1/4	65		15				20		<1		<1									
75	Do.		Not exposed.	Do.	Do.	1/8-1/4	55		20				25		<1		<1									
76	Do.		Do.	Lenticular	Do.	1/8-1/4	60		20				20		Trace		<1									
77	Do.		Do.	Do.	Do.	1/4	60		25				15		Trace		<1									
78	Do.		Do.	Do.	Do.	1	40	-	25				35		Trace		frace.							Beryl.	2 Crysta	10
79	Do.		Do.	Do.	Do.	1/4	59		15				25		1		Trace.						N. T.			
80	Do.		Crossoutting.	branching.	Do.	1/4	60		20				20		<1		Prace.							Beryl.	2 Crysta	10
81	Tonalite and fine-grained granite		Do.	Lenticular- branching.	Do.	1/4	62		15		5	raphic lbite	22		1		<1							Beryl.	1 Crysta	1
82	Do.	None,	Do.	enticular	Do.	1/8-1/4	75		5				20		<1		frace.			e militar						
83	Do.	Bone.	Do.	Do.	Wall sone Core	1/5-1/4	71		3				25 98		1		race.									
84	Tonalite.		Not exposed	Do.	One wait	3/4-1	45		25				30		Trace		Prace.							Beryl.	2 Crysta	10.
85	Tonalite and fine-grained granite.			Do.	Do.	1/16	62		3		5	raphie lbite	35						1							

humber and	Wall r	ock													Pe	gmatite	0									
number and name of peg- natite	Type and formation	Alteration	Relation to	Shape	Internal	Texture											м	ineral	ogy							
(Pl. II)	formation		Wall rock		structure	inches	Plagi	oclase	Pert	hite	Grap	hic ite	Qua	rtz	Muse	covite	Gas	rnet	Tour	maline	Lithiu	m minera	d.	Othe	r mineral	
							Per-	Size (inche	Per-	Size	Per-	Size nches	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Mineral		Size inches	Mineral	Per- cent	Size (inche
86	Conalite and fine grained granite.		Crosscutting	Lenticular branching	- One unit	1/4	55		20				25		< 1		Trace							Beryl. 2	crystals	
87	Tonalite.	None.	do.	Irregular	All the later of t	1/2-3/4	50		25				25		Trace		Trace								1000	
88	Hornblende gneis			Lenticular	-Northern	1/4	65		10				25		< 1		Trace							Beryl. 1	crystal.	1/4
		secret S		branching.	Southern	1/16	93		Trace.				7			Dig. 7-					e Nicelia			Epidote.	Trace.	-/
					end.																				List Car	
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	Wall r	ock													Pe	gmat1t	8									
humber and name of peg-	Type and	Alteration	Relation to	Shape		Texture											M	neral	реу							
atite (Pl.II)	formation		wall rock		structure	inches)		oclase	Pert	thite	Grap	hic	Qui	artz	Mus	covite	Gaz	net	Tour	naline	Lithium	n minere	ls	Othe	r mineral	.8
							Percent	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Si ze	Per-	Size	Mineral	Per- cent	Size	Mineral	Per- cent	Size (inche
89	Tonalite.		Not exposed.	Lenticular branching	- One unit	1/8	70		10				20		<1		Trace									
90	Coarse-grained granite.		do.	Lenticular		1/8	76		4				20		<1		Trace									
91	do.		do.	Irregular.	do.	1/8	73		2		e dina		25		<1		Trace									
92	do.		do.	Lenticular	. do.	1/2-3/4	60		20	3-7			20	1/4-2	<1		Trace.									
93	Tonalite.		do.	d.	do.	1/64	68		1				30		1	180	Trace.	1/8								
94	Hornblende gneise and fine grained granite	None.		do.	do.	2	20		45				35		Trace											
95	Tonalite and coarse-grained		Crosscutting.	do.	Northern end.	1/4	65		10				24		1										35.5	
	granite.				Southern part.	1	34		40				25	1-6	1											
96	Tonalite.	None.	do.	Lenticular branching	- Wall sone northeast part.		28		50		60-70	12-24	20		2											
			1		Wall zone southwest part.	1/4	66		12		2-3	6	20		2											
					Cores.	8	10		41				45		4	1/2-2										
97	Tonalite and fine grained granite.	⊢ do.	do.	do.	Wall zone southwest end.		15		63		80	4-12	20		2		Trace.									
					Wall sone northeast end.	A			15				30		1		1									
00				3012	Cores.	6-8	5		40				54		1									ALC: N	To lea	
98	Coarse-grained granite.	do.		Lenticular.	Wall zone Core.	8	177		10				15 68		3 2		Trace.							Biotite.	<1	
99	do.	do.	7 - 1 - 4	do.	Wall zone. Core.	1/4	162		20 30				18		77		Trace.							Biotite.	Trace.	
100	Tonalite and coarse-grained granite.	do.	Conformable.	do.	One unit.	1/8	79		1				20				Trace.									
101	do.	do.	crosscutting.	Lenticular-	do.	1/8-1/4	72		8				19		1		Trace.				1 A 80 11 13			Magnetite.	1 cryst	,
102	do.			branching.	West branc	1.1/4-1/	2 67		15				18		<1		frace.							Beryl.	100	. 1/4
103	Conslite.			Lenticular.	Manager and the	March March	67		15				18		<1		<1									
104	do.			do.	do.	1/16-	74	15.0	1				25				Trace.	1/8								
105	Tonalite and coarse-grained		do.	Lenticular- branching	Northeast and south-	1/32	74		1				25			The same	Trace.		4							
	granite.				est branch Morthward- trending branch.		60		20				20		Trace		Trace.							Beryl.	4 crysta	10. 1/2

mber and	Wall ro	ck					_	-		100					Pog	mat1 te										
me of pag-	Type and formation	Alteration	Relation to	Shape	Internal	Texture		dia.		100	2						M	neral	V30							
Pl. II)	Tornation		Sall Pock	-	structur(nches)	Plagic	clase	Pert	hite	Gray		Qua	rts	Muso	ovite	Gaz	rnet	Tours	naline	Li thius	ninere	10	Other	mineral	•
							Per- cent	Size	Per-	Size		Size inches	Per-	Size inches	Per- pen(Size	Per-	Si se	Per-	Size	Mineral	Per-	Size	Kineral	Per-	Size (inch
106	Tonalite.		Cresonting.	Lenticula branchin	r- One uni	. 1/4	72						20		<1		Trace.							Beryl.	2 crysta	10. 1/8
107	do.			40.	40.	1/8	67						25		trace.									Magnetite.	l crysta	
106	40.			Leaticular	. 40.	1/16-	n		3				25		trace.									Chlorite.	2	
109	do.			40.	40.	1/4	70		10			UNIE	20		<1		Trace									
110	60.		Not exposed,	Lenticular branching		1/6	71						25		<1		Trace.		Trace							
111	40.		do.	Leaticular	. do.	1/8-1/4	70		5				25	2001	<1		Trace						Falls			
112	Hornblende gneis	•	40.	Lenticular branching	4 0.	1/8-1/4	75		5				20		<1		<1							Beryl. 1	3 crysta	10. 1/
113	Hornblende gneiss and coarse-grain granite.	e4	40.	do.	do.	1/8-1/4	74		•				18		<1		Trace									
114	Hornblende gneis		do.	Lenticular	. 40.	1/8-1/4	70		10		1.00		20		<1		<1							Beryl.	2 crysta	10. 1/
115	Hornblende gneiss and granite.		40.	Lenticular branching		1/8	73		5				55		<1		Trace									
116	Hornblende gneles		Crosscutting.	Irregular	do.	1/8-1/4	65		15				20		<1		Trace.							Beryl.	2 crysta	10. 1/1 3/1
117	Fonalite and granite.		Not exposed.	Lenticular	40.	1/4-1/2	65		15				20		<1		Trace.							do.	3 cryste	10. 1/
118	Coarse-grained granite.	Hone.		Lenticular branching	do.	1/16	81		*				15		<1		Trace.									
119	do.		Not exposed.	do.	do.	1/8-1/4	67		8				25		race.		Trace.	***								
120	fonalite.		Crosscutting.	do.	do.	1/32-	50		2				18		trace.		Trace.									
121	do.		do.	do.	do.	1/16-1/0	76		*		*		20		race.		Trace.							Beryl.	5 crysta	10. 1
122	Not exposed.		Not exposed.	Irregular.	do.	1/8-1/4	65		15		5		20		<1		Trade.				V 8 6					
123	do.		do.	do.	40.	1/8-1/4	62		18				20		<1		≺1						C date			
124	Fine-grained granite.		do.	Lenticular branching	do.	1/8-1/4	70		10		2		20		<1		Trace.									
125	Fonalite.			Lenticular.	do.	1/8	75	1	5	5,334			20		race.		Trace.	7	E850					Beryl.	1 crysta	1. 5/
126	do.		Not exposed.		do.	1/8-1/4	75		10		100		15	-	race.		Trace.						7616	do.	1 crysta	1. 1/
127	fornblende gneise		do.		do.	1/8-1/4	70		10				20		4		Trace.					1		- 1 V-0		
126	do.		do.		do.	1/8-1/4	73	Oto 3	7				20		<1	24.4	Trace.		1					Beryl.	2 crystal	10. 1/
129	Tonalite.		Crosscutting.	Lenticulen	40.	1/8	76		4	7		The state of	20	11.3	race.	Page 1	Trace.		1	1	Beer to be to	123 4	1	do.	l crystal	G. J.

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unber and	Wall r	ock				1000									Pog	nat1 te	•			White.		100				
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture								4			361	neralo	er .							
(P1.II)	Torraction		ANTI LOCK		structure	120000)	Plagi	oclase	Pert	hite	Grapi	hic ite	Quar	ts	Muson	ovite	Gar	net	Tours	aline	Li thius	miner	ı.	Othe	r mineral	
							Percent	Sise	Per-	Size	Per-)cent(Size inches)	Per- cent(1	Size nches	Per-	Size	Per-	Si se inches	Per-	Size	Mineral	Per-	Sise	Mineral	Per-	Sis (inche
130	Tonalite.		Grosscut ting.	Lenticular branching	- One unit	1/8-1/4	75		5				20		<1		Trace.							Beryl.	3 crystal	o. 1/0
131	do.		40.	40.	40.	1/4	65						27		Trace.		Trace.	1 8					7.	40.	2 orystal	- 1/0
132	40.	1.1	40.	40.	do.	1/8	77		5	100			18		<1		Trace.								2 crystal	
133	Hornblende gnete		Not exposed.	do.	do.	1/4-1/2	56		4				40		<1		Trace.							40.	2 orystal	
134	do.		40.	Irregular.	do.	1/8-1/4	73		2				25		a		Trace.							do.	2 crystal	Barrie
135	do.		do.	Lenticular	. do.	1/8-1/4	70		10				20		<1		Trace.							40.	1 crystal	
136	Tonalite.	None.		Irregular.	do.	1/2-3/4	70		8				22	1	<1		Trace.							40. 1	l orystal	
137	Hornblende gneis			Lenticular branching		1/4	75		10				15		4		Trace.							40.	a crystal	. 1/1
136	Hornblende gneis	None.		do.	40.	1/8	72		8				20		<1		Trace.									
139	Hornblende gneis	. do.		Lenticular	do.	1/4	65		15				20		<1		Trace.		Trace					Beryl.	1 crystal	1/9
140	Coarse-grained granite.	do .		Lenticular branching	40.	1/2-3/4	60		30				10		<1	4.50	Trace.						1			7-
141	Tonalite and granite.	do.	Crossoutting.	do.	do.	1/4	72		30				15				Trace.									
142	remalite.			Lonticular,	do.	1/4	72	1	3				25				Trace.									
143	40.			Irregular.	do.	1/8	73		2			4	15		race.		Trace.							Beryl.	l orystal	3/1
144	Hornblende gneise		Not exposed.	Lenticular	do.	1/2	40		35			907	80												-1	-1
145	do.		do.	do.	40.	1/16-	81		4				15				Trace.								1 1	
146		Bone.	Grossout ting.	Lenticular branching.	do.	1/2	55		25				80		Q		Trace.									
147	Tonalite.	None.		Lenticular	do.	1/4	70		race.				30											Beryl.	3 crystal	. 1/1
146	40.	do.		do.	do.	1/8-1/4	73		2				25		<1		Trace.							40.	l orystal	
149	40.			do.	do.	1/8	72		3		1		25				Trace.							100		
150	fonalite and coarse-grained granite.	Sone.		lenticular- branching.	do .	/8-1/4	65		15			ľ	20		race.		<1							Beryl.	2 crystal	. 1/
151	40.			enticular.	do.	1/4	70		15			,	15													
152	Coarse-grained granite.		Bot exposed.	do.	do.	1/8	70		15			,	15				frace.									
153	Coarse-grained granite and horn blende gneiss.	Scae.	Grossoutting.	enticular- branching.	do.	1/4-1/2	63		15				12		(1		frace.									

	Wall r	ock						9							Pegn	atite										
umber and ame of peg-	Type and	Alteration	Relation to	Shape	Internal												Mir	eralo	gy						AST TO	
atite (Pl. II)	formation		wall rock		structure(inches)	Plagic	clase	Pert	hite	Grapi	hic ite	Quart	tz	Musco	vite	Garı	et	Tour	aline	Lithiu	m minere	ls	Othe	r mineral	8
							Per-	Size	Per-	Size inches	Per-	Size inches)	Per- S	Size nches)	Per- cent(1	Size nches	Per-	Size nches	Per-	Size	Mineral	Per-	Size	Mineral	Per- cent	Size (inches)
154	Not exposed.		Not exposed.	Irregular.	One unit.	1/8	73		7				20		<1		Trace.									
155	Hornblende gneis and coarse- grained granite		do.	Lenticular	do.	1/8	72		8				20	T:	race.											
156	do.		do.	do.	do.	1/16-	80		4				16	•	race.									Biotite.	Trace.	
157	Tonalite.		do.	do.	do.	1/8	70		10				20	r	race.		Trace.								1.000	
158	Hornblende gheis			do.	do.	1/2	55		20				25	r	race.		Trace.									
159	do.		Not exposed.	do.	do.	1/2	55		20				25				Trace.									
160	Coarse-grained granite.			do.	do.	1/16-	80		5				15	1	race.		Frace.							Martite.	3 crystal	
161	do.	None.		Lenticular branching		1/8	75		5				20	T	race.		<1									
162	Tonalite and coarse-grained granite.	do.		do.	do.	1/8-1/4	70		10				20		race.		<1							Beryl. 8	crystals	1/8-
163	do.		Crosscutting.	do.	do.	1/4-1/2	65		20				15				<1							Beryl. 1 Biotite. Martite.	crystal. Trace.	1/8
164	Coarse-grained granite.		Not exposed.	do.	do.	1/8-1/4	60		25				15	**	race.		race.							Martite.	Trace.	
165	Tonalite and coarse-grained granite.	None.		Lenticular.	do.	1/8-1/4	70		10				20		<1											
166	do.	do.		do.	do.	/8-1/4	65		15				20		<1											
167	Tonalite.		Not exposed.	Irregular.	do.	1/4-1/2	60		20				20			1	race.									
168	Hornblende gneiss and granite.	Mone.		Lenticular.	do.	1/32- 1/16	75		1				25													
169	Coarse-grained granite.	do.		Lenticular- branching.	do.	1/4-1/2	63		25				12		<1		<1									
170	Tonalite and granite.	do.		Lenticular.	do.	1/4	60		15				25		<1	•	race.									
171	Hornblende gneiss		Not exposed.	Lenticular- branching.	do.	1/8	65		20				15				<1									
172	do.		do.	Lenticular.	do.	1/8-1/4	65		12				23			+	race.							Beryl. Martite.	crystal	. 1/8-1/
173	do.		do.	do.	do.	1/8	73		7				20	T	race.	+	race.							Beryl.		. 1/8-
174	Hornblende gneiss and granite.	None.		Lenticular- branching.	Wall zone. Core.	1/8-1/4			10			1	10	T	race.	+	race.							Martite. Beryl. 1	Trace.	

175

umber and	Wall z	roak						-		10012				Pe	gmat1 to	•					-				7014
ame of pag-	Type and formation	Alteration	Relation to	Shape	Internal	Texture										365	neralo	EY							
atite (Pl. II)	IOFMECTOR		wall rock		structure	Indias)	Plagic	clase	Pert	hite	Graphic granite	Qu	arts	Mus	covite	Gar	net	Tours	aline	Lithium	miners	de .	Other	mineral	
							Per-	Size	Per-	Size nches	Per- Sis	e Per-	Size	Per-	Size	Per-	Si se Inches	Per-	Size	Mineral	Per-	Size	Mineral	Per- cent	Si (incl
175	Coarse-grained granite.	Hone.		Lenticular branching		1/4	65		15			20		Trace											
176	Fonalite and coarse-grained granite.	do.		Lenticular	Wall sone	1/16-	83 10		5			15		Trace											
177	fonalite and granite.		Not exposed.	Leaticular.	One unit.	1/8-1/4	65		15			20				<1									
176	Coarse-grained granite.	None.		do.	do.	1/8	75		5			20				frace.							Martite.	Trace.	
179	Tonalite.		Not exposed.	40.	do.	1/8	73		7			20				Prace.							Magnetite.	Trace.	
160	do.			Lenticular- branching.	do.	1/2	45		25			30				Prace.									
161	fonalite and granite.	Bone.	Crosscutting	. do.	do.	1/4	72		20			8		Trace		frace.							Magnetite	Trace.	
162	do.	40	do.	Lonticular.	do.	1/16	66		25					Trace		<1									
183	do.	do.	do.	do.	do.	1/16-	70		5			20	l'	Trace		<1								*	
184	40.		Not exposed.	do.	do.	1/8	70		20			10				Prace.							Beryl. 2	crystal	. :
185	do.	None.	Crosscutting	do.	do.	1/2	65		25			15		Trace		race.									
186	do.	do.	do.	enticular- branching.	do.	1/4	71		20			8		1		frace.							Martite. Beryl. 1		,
187	40.	do.		enticular.	40.	1/4-1/2	50		32			18		1		<1									
188	do.	do.	Crosscutting	Lenticula branching.	r- do.	1/8-1/4	73		22			5				trace.							Magnetite. Beryl. 2		. :
189	do.	do.	do.	do.	do.	1/4-1/2	45		30			25				Trace									
190	Tonalite.	do.	do.	do.	do.	1/16-	69		6			25											Beryl. 2	crystal	
191	do.		do.	do.	40.	1/16	75		5			20				Trace									
192	Tonalite and granite.	Hone.		Lenticular	. Wall sone	. 1/32	83 3		2 37			15 60													
193	Tonalite.	do.		Lenticular branching	- One unit.	1/16	76		4			80				<1							Magnetite,	Trace.	
194	do.	40.		Lenticular	. do.	1/16	78		2			20				Trace							do.	Trace.	
195	Tonalite and granite.		Not exposed.	Lenticular branching	- do.	1	33		45			55		<1		<1									
196	do.			do.	do.	1/4	70		10			20		Trace		Trace					25.73		Beryl. 2	crystal	8.
197	Tonalite.	None.		Lenticular	do.	1/8	67		8			25		<1					1				do. 5	crystale	1

	Wall r	ock													Peg	matite						4949				
Aumber and	Type and	Alteration	Relation to	Shape	Internal	Texture											16	neral	E7							
(Pl.II)	formation		wall rock		structure	inches)	Plagi	oclase	Per	thite	Gray	phic nite	Qua	arts	Musc	ovite	Gaz	net	Tours	aline	Li thius	miner	10	Other	r mineral	•
							Per-	Sise		Sise	Per-	Sise								Size	Mineral	Per-		Mineral	Per-	Sise (inches)
198	Tonalite.		Not exposed.	Lenticular	. One unit	1/4-1/2	50		35				15		<1									Beryl, 1	orystal	. 1/4
199	Fine-grained granite.			Oval.	do.	1/2-3/4	66						30		<1											
200	40.			Irregular.	Wall sone	. 1/8	80 15		10				10		<1 <1											
201	do.			Lenticular	. One unit.	1/8-1/4	17						15		<1		Trace							Beryl. 1	orystal	1/2
505	Hornblende gnois and fine-grains granite.	None.		Irregular.	40.	1/2-3/4	74		5				20		1											
203	Pine-grained granite.			do.	do.	3/4	58		20				20		2											
204	do.		Not expessed.	Lenticular	. do.	1/2	45		35			1000	20		<1		Trace.								1000	
205	Hernblende gneis	. None.		Irregular.	Wall sone. Core.	1/8	77 20		30				a0 50		<1 1		Trace.	,	cryst	al.				Gahnite.	crysta	
206	40.		Not exposed.	Lenticular	. One unit	3/4	balls.		25				30		1									tantalite	2 cryst	d
207	Tonalite.		do.	do.	40.	1/4	59		7				30				Trace.									
206	Pine-grained granite.		do.	do.	Vall sone	1/4	162		5 15				30		3											
209 Opportunity Dike No. 10			Crossoutting.	Lenticular branching		1/8-1/1	60		20				20		frace.											
210					southern bo	ly. 2	170		10				20		<1											
Opportunity Dike No. 9	40.		Not exposed.	Lenticular	One unit.	1	158		80				20		2				1							
211 Opportunity Dike No. 7	do.		40.	do.	do.	1/8	78		7				15		Trace.		Trace.									
212 Opportunity Dike No. 6	do.		do.	do.	40.	1/8-1/4	714		5				20		Trace.		Trace.		1							
213 Opportunity Dike No. 5	do.		do.	do.	Wall some Core.	16	170 120		12 40				18		<1 <1				<1					Beryl.	Several.	Up to 3
214 Opportunity	do.		do.	do.	North and	3 - 5	115		2				81		2											
Dike No. 4.					Central section.	3/4	65		15				20		<1				Trace.							
215 Opportunity Dike No. 3	do.		do.	Lenticular branching	- Small . stringer albite-qua- perthite pegmatite.	is- ^{1/4}	65		15				20		< 1									Beryl. 1	erystal.	1/4

	Wall r	ook				737	MIST.	Die No.		ion is					Pog	mat1t										
Number and name of pag- matite	Type and formation	Alteration	Relation to	Shape	Internal	Texture											16	neral	ey .							
(F1.II)	1011111111		Wall Four		***************************************	1	Plagi	oclase	Pert	hite	Gran		Que	urts	Musc	ovite	Gaz	rnet	Tours	aline	Lithium	miner	als	Other	mineral	•
							Percent	Size	Per-	Sise	Per-	Sise	Per-	Size	Per-	Sise	Per-	Sise	Per-	Size	Mineral	Per-	Size (inches)	Mineral	Per-	Sise (inches)
215 Opportunity Dike No. 3 (Continued)	Fine-grained granite and hornblende gnei	.	Not exposed.	branching	- Central . part. Cleavelan- dite-quart pegmatite.	1	169						80		2		Trace	Up to	. 1	1/4-1	Lepidolite	<1		Topas. Monasite. Columbite- tantalite.	crystal	s. 1/2
					South-centunit. Perthite- cleavelendi quarts pagmatite.	2 - 3	¹ 20		60				20		<1		Trace.		a					Microlite. Beryl. 20	crystal	1/32-1/4 0.3/4-1
					Pod. Quarts-albi pagmatite	to 3 - 4	20		2				75		3				<1				La 4			
216 Opportunity Dike No. 2	do.		do.	do.	North- trending branches.	1/8-1/4	72		10				18		<1											
					Fortheast branch,	1-1/2	173		5				20		2		<1							Microlite. Beryl. 5	2 orystal	10. 1/4
					Central and southweste branch.		1g		64	12-36			25		3	1/2-3	•								crystal	
217 Opportunity Dike No. 1	Pine-grained granite.		40.	Oval.	One unit.	4 - 8	5		10				gh.		1				< 1							
Opportunity	Pine-grained granite and hornblende gnei	6.	do.	Lenticular.	do.	2	155		30				15		<1				Trace.							
	coarse-grained granite.	Fone.		Irregular.	do.	1/8	75						20											Biotite. Chlorite.	<1	
220	do.			Lenticular,	do.	1/4	75		3		100	I STATE	20					NO.	7-11-1					do.	2	
221	ine-grained granite.		Not exposed.	do.	do.	1/4	65		15				20		<1		Prace.									
222	overed.		do.	Lenticular- branching.	do.	1/8-1/4	56		5				35											Chlorite.	h.	
223 1	ine-grained gneissic granit	None.	Conformable.	do.	do.	1/4	55		25		8		80		<1		Trace.							Beryl. 1	orystal.	1/4
224	do.	do.	do.	Lenticular.	do.	1/8	77		3				80				Trace.	J. S. S.		E. P.						
225 1	ine-grained granite.		Not exposed.	do.	do.	1/8	72						20				Trace.		Trace.					Magnetite.	Trace.	
	ine-grained granite and hernblende gneis			40.	do.	1/4	65		15				80		Trace		Trace.		Trace.					Beryl. 5	orystal (. Up to

670 S FS0

Number and name of peg- matite (F1.II)	Vall r	ock		COOK IN				e au a						Gilli	Pegmati	te					The same				
	Type and	Alteration	Relation to	Shape	Internal structure	Texture (inches)																			
	formation		wall rock				Plagioclase		Perthite		Graphic granite		Quarts		Muscovite		Garnet		rmaline	Lithiu	Lithium minerals		Other minerals		le .
							Per-		Per-	Size inches	Per-	Size inches	Per- S	ize l	Per- Sis	e Pe	er- Sise	Per		Mineral	Per-		Mineral	Per-	Size (inche
227	Fine-grained granite and herablende gnei	Hone.	Conformable.	Lenticular	. One unit	1/4	70		10				20	1	race.	Tre	ace.						Biotite.	Trace.	
226	do.		Not exposed.	do.	do.	1/8-1/4	70		10	58			20			Tr	ace.								
229	Fine-grained granite and tonalite.		do.	Lenticular branching		1/8	77		8				15		<1	Tre	ace.	Tre					Martite.	Trace.	
230	Tonalite.			do.	Wall some. Gore.	1/4	70 10		10 30				20 60		Trace.	<	1						Beryl. 6	crystal	1/2
231	Tonalite and granite.		Not exposed.	Lenticular	.Wall some.	1/8-1/4	80 10		5 40				15		<1	<	1	<:					Beryl. 3	crystale	+
535	Tonalite.			Oval.	One unit.	1/8	74		3				22		<1	1	1/16								
233	Fine-grained granite.		Not exposed.	Lenticular	. do.	1/8-1/4	70		10				20		<1	Tre	ace.						Beryl.	crystal	1- 1/4
234	Hornblende gneiss.		do.	40.	do.	1/4	69		5				25		1	Tre	ace.								
235	Fine-grained granite.		do.	do.	do.	1/4	60		15				25		Trace.										
236	Fine and coarse- grained granite		Conformable.	do.	do.	1/4	65		15				20		Trace.								Biotite.	trace.	
237	Fine-grained granite.	do.		Lenticular branching		1/4	60		20				20		<1	Tre							Beryl, 2	prystale	3/8
238	do.	do.		Lenticular	do.	1/2-3/4	54		20				25		1	Tre	sce.	Tree					do. 3	prystale	1/4
239	Hornblende gneis		Not exposed.	Oval.	do.	1/8-1/4	70		10				20		Trace.	Tre	sce.						do. 1	prystal.	. 1/4
240	do.		do.	Lenticular	do.	1/4	70		15				15		Trace.	Tre	ace.						do. 1	prystal.	. 1/2
241	Gneissic granite	1	do.	Oval.	do.	1/2-3/4	41		40				18		< 1	Tre		Trac		6-21					
545	Fine-grained granite.	Fone.		do.	do.	1/2	50		30				20		1	Tre	100.								
243	do.	do.	Crosscutting.	Lenticular branching	do.	1/4	70		10				20			Tre	LOO.						Martite.	frace.	
5##	do.	do.		Lenticular	do.	1/4	65		10				25		<1	Tre	LCO.						Beryl. 4	prystals	1/8-
245	Hornblende gneis tonalite and granite.	, do.	Crosscutting.		Wall zone.	g	57		20		25 2	4-36	20		3	-	.ce.						do. 1	crystal.	3/4 by
	granite.				Wall zone.	1/8	85		3				12		<1	Tre	ice.						do. 2 Columbite- tantalite		
					Core pods.	1 - 2	49		20	NO.			30		1	100			100	1				,,	

Quartz

Per Size Per Size Per Size Per Size Per Size Per Size Per-

cent (inches)cent (inches)cent (inches)cent (inches)cent (inches)cent (inches)cent

12

25

25

65

20

35

15

15

15

35

25

20

25

25

20

12

20

20

10

25

Graphic

grani te

36 10-19 12-36

Pegmatite

Muscovite

<1

<1

3

2

<1

<1

<1

<1

<1

<1

<1

Trace.

<1

<1

<1

<1

<1

×1

Mineralogy

Tournaline

Size

Lithium minerals

Per- Size

cent

Mineral

Other minerals

Per

cent

Size

inches

Mineral

Garnet

Trace.

Trace

Trace.

Trace.

Trace

Trace

Trace

Trace

Trace

Trace

Trace

Trace.

Trace

<1

Trace.

1/4-1/2 Trace.

Wall rock

Hornblende gneise. None.

Alteration

Hone.

Relation to

wall rock

Crosscutting Irregular Wall sone

Lenticular.

branchine.

do.

do.

Lenticular.

Lenticular-

branching.

Benticular.

Lenticular-

branching.

rregular.

Lenticular,

Irregular.

enticular

Crosscutting Lenticular

do.

Not exposed.

do.

Not exposed.

Internal

Core pode

One unit.

Wall sone.

Core.

One unit.

do.

do.

do.

do.

do.

do.

do.

do.

do.

Wall sone.

Core.

Core pode. 6 - 8

Wall sone. 1/4-1/2

Texture structure (inches)

1/2

3/4

1-1/2

1/8-1/4

/8-1/4

1/4

1/4

1/4

/8-1/4

1/4

1/4

1/4

1/4

/8-1/4

1/2

1/4

/16-1/8

Plagioclase

65

64

53

70

80

75

70

57

60

76

69

53

55

55

73

72

33 1/4-1/2

Perthite

35

40

30

10

51

10

15

15

10

22

20

25

15

30

5

Type and

formation

do.

de.

do.

Tonalite.

do.

Tonalite.

do.

granite.

granite.

Fine-grained

Fine-grained

granite.

do.

Hornblende gneise and fine-grained

Hornblende gneise.

Hornblende gneiss.

Hornblende gneiss.

Hornblende gneiss.

Mumber and

(Pl. II)

246

249

250

251

252

253

254

255

256

257

259

260

261

262

263

264

265

Opportunity

2A and 3A

matite

name of peg-

Beryl. 1 crystal. 2-1/2 Monarite.1 crystal. 1/8 Biotite. 2 crystal. 6 by 10 Beryl. 2 crystals.1/4-1 do. 4 crystals.1/5-1-1/2 constite. 2 crystals. 1/5-1-1/2 columbite- tantalite.2 crystals. 1/32- 1/4 Biotite. Frace.				
	Biotite. Beryl. 1 Beryl. 2 do. 4 Monasite.5 Columbite- tantalite.	<1 crystal crystal crystal crystal 2 crysta	. 6 by 10 e.1/4-1 e.1/5-1-1 e. Up to 1	/2
Beryl. [14 crystals. 1/8-1/2 crystal. 3/4	Biotite.	Trace.		(
	Beryl. [14]	crystals crystal.	. 1/8-1/2 3/4	

400 0 750

mber and	Wall r	ock										e per			Peg	mat1te						Pale	10000			
me of peg-	Type and	Alteration	Relation to	Shape	Internal	Texture											Min	eralo	gy .							
tite Pl. II)	formation		wall rock		structure	(inches)	Plagic	oclase	Pert	hite	Graph	nic i te	Qua	rtz	Muso	ovite	Garn	et	Tourn	aline	Lithium	miners	ı.	Other	mineral	•
							Percent	Size	Per-	Size nches	Per-	Size nches)	Per-	Size inches)	Per-	Size	Per-	Size iches)	Per-	Size inches	Mineral	Per-	Size	Mineral	Per- cent	Siz (inche
266	Pine-grained granite.		Not exposed.	Lenticule	.Vall sone Core.	1/4	70 45		10 25				20		<1 <1		Trace.							Columbite- tantalite	. 1 cryst	al.
267	Hornblende gneie	•	do.	Oval.	Wall sone. Core.	1/8-1/4	69 24		3 30				25 45		3		Trace.									
268	Hornblende gneis and fine-graine granite.	ė.	do.	Lenticula	. One unit	1/4	67		3				30		<1											
269	Fine-grained granite.		do.	do.	do.	1/4	61		4				35		<1											
270	Hornblende gneis	s. None.		Irregular.	Hanging wall.	3/4-1	111		25				30		1											
	*				layer. Hiddle layer.	6 1/16	25		30				45 15		<1	1-2	<1							Beryl. 1	crystal.	1 by
					layer.																					
271	Tonalite.	do.		Lenticular	. One unit	3/4	40		35				25		<1		Trace.									
272	Hornblende gneis	s. do.		Irregular.	do.	1/8	80		5				15		<1		<1									
273	do.	do.		do.	Wall sone. Core pode.	1/8	76 40		8 25				15 35		1 <1		Trace.									
274	do.		Crosscutting.	Lenticular branching	- Hanging wall layer.	3/4	45	1/8-1/	4 30	1-4			25		<1		Trace.							Martite.	Trace.	1/2
					Foot wall layer.	1/16	81		3				15		1		<1									
					Core.	2	25		40				35		<1								. *	1		
275	do.				Well sone. Core pods.	1/8-1/4	58	,	3 20				35 35		1											
276	do.	None.		Lenticular branching	One unit.	1/4	74		5				20		1											
277	Hornblende gneis and granite.	do.	Crosscutting.	Irregular	. Well sone	.1/8-1/1	71 30		8 54				20 15		1											
278	do.	do.		Lenticular	. One unit.	1/4	64		7				27		2		Trace.									
279	Hornblende gneis	•		Lenticular branching	Wall sone	. 1/4	65		15				20		<1		Trace.							Biotite.	Trace.	1/8-
					Core.	3	30	1/8-1/4	30				39		1	1/4-3/	4 <1							Beryl. 35 Gahnite. Samarskit	Trace.	1/64 ystal
280	do.			Lenticular	One unit.	1/8	73		12		1		15		<1		Trace.							Martite.	Trace.	
281	do.			do.	do.	1/8	75		10				15		Trace		Trace.	6		20691				Beryl. 3	crystals	.1/8-
282	do.			Lenticular branching		1/4	65		20				15		Trace		Trace.							Biotite.	Trace.	
283	do.		Not exposed.	Lenticular	do.	1/2-3/4	60		15	3 - 6		150.00	25		<1		<1		48.34	NO F		Tay (ch)	9.45	Beryl. 1	crystal	1/4

Number and	Well r	ock									- 14				Peg	matite		B								
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture	Bes										Mine	ralo	gy							
(P1. II)	lormation		Wall rock		structure	inches		clase	Pert	hite	Gran		Qua	rtz	Musc	covite	Garne	t.	Tour	aline	Lithium	miner	ds	Other	mineral	•
							Per- cent	Size (inches	Per-	Size	Per-	Size	Per-	Size (inches	Per-	Size (inches	Per- S bent (1)	ize	Per-	Size inches	Mineral	Per-	Size	Mineral	Per-	(inche
284	Fine-grained granite.	None.		Lenticular	.One unit.	1/8	70		8				22				Trace.									
285	do.	do.	Conformable.	do.	do.	1/8	70		15				15				<1	-								
286	do.			do.	do.	1/2	55		25				50		<1		Trace.									
287	Hornblende gnei	•.		Lenticular		1/4	75		15				10		Tras		Trace.									
288	do.		Not exposed.	Lenticular	.Wall zone Core.	1/4	160		<1				30 40		4 <1		<1	1	cryst	al.				Beryl. 3	crystale	. 1
289	do.		do.	do.	Wall zone. Core.	1/2	160		2 5 5	2 - 3			30 35		3 <1	X	Trace		crysts crysts					Beryl. 1 Columbite-		
290	Hornblende gneis and granite.	None.		Lenticular branching		1/5	79		2				15		4		Trace.								Trace.	
					Center layer.	4 - 6	20		50				30		<1		Trace.		Thace.					Beryl. 1	crystal.	
					Pootwall layer.	3/4	37		42				20		1		Trace.							Biotite. Columbite- tantalite Monazite.	. 1 crys	tal.
291	Fine-grained granite.		Crosscutting.	Lenticular	. One unit.	1/2	55		20				25		<1		Trace.									
292	Tonalite.		do.	Lenticular branching	. wall	3/4	54		25				50		1											
					layer. Pootwall layer.	1/16	86		3				10		1											
293	do.			Oval.	Core.	1/8	78 15		5 5				20		<1 1											
294	Hornblende gneis	•		Irregular.	One unit.	1/8	66		3				30		1								1			
295	do.	None.		do.	do.	1/4	70		10				20		<1		Frace.									
296	do.		Not exposed.	Lenticular.	North and south ends Center unit.	1/16	25 83		43				30 15		2		<1									
297	Hornblende gneis	None.	Crosscutting.	Irregular.	Wall sone. Core.	1/2	74 14		5 43				20		1 <1	-										
298	Hornblende gneis		Not exposed.	Lenticular	One unit.	1/2	65		10				25		Trace											
299	Covered.		do.	Lenticular branching.	do.	3/4	65		20				25		<1									Beryl. 8	crystal	. 1/1 1/4
300	Hornblende gneis		do.	do.	do.	3/4	53		22	3-6			25		<1		Trace.								crystal	.1/4
301	do.		do.	Irregular.	do.	1/4	66		3				30	128	1	14	<1	186						Beryl. 2		

lumber and	Wall r						0.000		1						? Pog	natite	-1									
ame of pag-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture	detta										165.1	eralo	EY							
(M.II)							Plagic	clase	Pert	hite	Grap	hic ite	Quar	ts	Muso	ovite	Garı	net	Tours	aline	Li thium	minere	de .	Other	mineral	•
							Per-	Size	Per-	Size	Per-	Sise	Per-	Size	Per-	Size inches	Per-	Sise	Per-	Size inches	Mineral	Per-	Sise	Mineral	Per-	(inche
302	Hornblende gneie	•	Not exposed.	Lenticular branching	- Wall sone	1-1/2	70 15		10				20 45		<1 <1		Trace.							Beryl. 2	2 orystal	0.1/2-1
303	do.		do.	Lenticular	. One unit.	1/8	80		8				12		<1		<1									
304	do.		do.	Oval.	do.	1/2	54		20				25		1	300	<1								1	
305	Hornblende gneis and fine-grains granite.			Irregular.	Wall some.	1/8	80		5		1		15		<1		<1							Beryl. 3 Chlorite.	crystal Trace.	. 1/1
306	Hornblende gneis		Not expessed.	Lenticular		1/4	68		10	3			22													
Opportunity No. 4					West layer	1-1/3	179	1-3	10	,			15		3		Trace.		<1 <1		Lepidolite	4-5	<1/16-	Microlite.	2 cryst	10. 1/
307 Opportunity No. 4	do.		do.	Irregular.	One unit.	1/2	47		30		5-10		20		3		<1		Trace.							
308	do.		do.	Lenticular	Wall sone. Core.	1/2	59		15		10-20		25		1		trace.							Beryl. 1	crystal	
309	do.		do.	Irregular.i	layer.		64		20				15		1		<1									
					layer.	4-6	5		75		>50		80		<1											
310	do.		do.	do.	One unit.	1/4	35		45		5-10		20		1											
311	do.		do.	Lenticular.	do.	3/4	fff		35		30	10000	20		1		<1		Prace.	7				Beryl. 2	erystals	1/4
312	do.		do.	Irregular.	do.	8-3	25		54		75		20		1		race.		frace.							
313	Granite gneiss.		do.	Dval.	do.		15		64		75		50		1											
314	Hornblende gneiss and granite gnei		40.	Lenticular.	do.	1	38		140		50		50		2											
315	Conalite and gran	ite. Sone.	Crossoutting.	Lenticular- branching.	do.	1/2-3/4	75		5				20		<1											
316	Conalite.	do.	do.	Irregular.	Core.	1/4	66		5 25		5		25		1	2	4							Beryl. 1	crystal.	3/4
317	do.			branching.	Core.	4	15		15 35				25		5		race.									
318	do.		do.	Lenticular.		1/4	69		1				30		<1											
319	Bornblende gneise		do.	40.	do.	3/4	Jele		30				20		6							Haling				
320	do.		do.	branching.	do.	1/2	61		1				35		3											
321	Tonalite.	None.	onformable.	Lenticular.	Wall sone Intermed- iate sone.	1/8	70 g4		2				30		4											
	A 44 4 7 8 9				Core.		S S S						00		1	1	-	3000				1789		2 - 1 - 5 / 7		

Number and	Wall ro	ock													Peg	matite										
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal												M	ineralo	EX							
(Pl.II)	Iormation		wall rock		structure	inches)	Plagioc	lase	Perth	ite	Graph:		Quar	rts	Musc	ovite	Gaz	rnet	Tour	naline	Li thiu	n miner	al.	Other	mineral	•
							Per-	Size	Per- cent(1	Size	Per-	Size Pe	nt(Size inches	Per-	Size inches	Per-	Size inches	Per-	Size (inches	Mineral	Per-	Size (inches	Mineral	Per- cent	Size (inche
322	Hornblende gneis		Not exposed.	Lenticular	. One unit.	1/4-1/2	62		10				25		3											
323	Tonalite.	Mone.	Crosscutting.	Oval.	Wall zone. Core.	1/2-3/4	69		10				20		1											
324	do.	do.	do.	Irregular.	One unit.	1/4	68		10				20		2											
325	do.	do.	do.	do.	do.	1/8-1/4	84		10	1			5	1/4-1/	2 1-2	1/2	<1							Beryl. 1	prystal.	
326	do.		Not exposed.	Oval.	do.	1	40		40				20		<1							702 (N				
327	do.		do.	do.	do.	1/2	52		30	6		1	8		<1											
328	do.		do.	Lenticular	. do.	1/8	80		5	2		1	5													
329	do.		do.	Lenticular branching		1/4	54		30	6		,	5		1		<1									
330	Tonalite and fine grained granite.		do.	do.	40.	1/4	60		19				0		1		Trace.							Beryl. 3 Biotite. Magnetite.	Trace.	dales.
331	Tonalite.		do.	Lenticular	do.	3/4	45		35			2	5		<1										PAUL S	
332	Fine-grained granite.			Lenticular branching		1/2	35		40			2	5		Trace		Prace.									
333	Tonalite.		Not exposed.	Lenticular	do.	1/4-1/2	50		30			2	0		<1		<1							Magnetite.	41	
334	Hornblende gniese and granite.	None.	Crosscutting.	Irregular.	do.	1/16	73		7			a	0		Trace		Prace.							do.	Trace.	
335	Hornblende gneiss		Not exposed.	Lenticular	do.	1/8	75		10			1	5				frace.									
336	do.		do.	do.	do.	1/4	55		25			2	0		<1		Trace.							Beryl. 1	crystal	1/8
337	Hornblende gneiss and granite.	None.	Crosscutting.	do.	do.	1/4	60		25			1	5		< 1	_	<1									
338	Tonalite.		Not exposed.	Lenticular branching.	do.	1/4-1/2	55		20			2	5		<1		frace.							Beryl. 2	crystals	1/4
339	do.		do.	do.	do.	1/4-1/2	55		25			2	0		<1		Trace.							Epidote.	Trace.	
340	do.		do.	Lenticular.	Wall sone. Core.	1/4	67		10 5				0		3		<1 1									
341	do.		Crosscutting.	Lenticular- branching.	One unit.	1/8-1/4	63		15			2	2				<1									
342	do.		Not exposed.	Lenticular.	do.	/8-1/4	70		10			2	0		<1		race.				7			Beryl. 2	crystals	1/4
343	do.		do.	do.	do.	1/16	78		2			2	0				race.									
3144	Conslite and granite.		do.	do.	do.	/8-1/4	55		25			5	0		Trace		race.									
345	do.	None.	crosscutting.	Lenticular branching		/8-1/4	65		15			5	0		<1		race.									
346	do.	do.		Lenticular.	do.	1/8-1/4	65		15			2	0		<1		race.					1				- 366

Aumber and	Wall ro	ock	12.60					90.00			2015			Pe	gmat1te	•	250		910	0.510-6					
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture										M	neralo	EX							
(Pl. II)	rormation		ANTI LOCK		Structure		Plagio	clase	Perthi	te	Graphi grani t	c Q	uartz	Muse	covite	Gar	net	Tours	naline	Lithium	miner	als.	Other	mineral	•
							Per-	Size inches)	Per- S	Size	Per- S	Size Per-	Size	Per-	Size	Per-	Si ze inches	Per-	Size (inche	Mineral	Per-		Mineral	Per- cent	(inches
347	Tonalite.		Not exposed.	Lenticular branching	one unit.	1/4	60		20			20		Trace		Trace.				•			Beryl. 2	crystal	
348	do.		do.	do.	do.	1/4-1/2	50		25			25				frace							do. 1	crystal.	1/4
349	do.		do.	Lenticular	. do.	1/2	40		25			35				Trace		W 3							
350	Tonalite and granite.		do.	do.	do.	1/8-1/4	65		15			50		Trac	•	Trace.							Magnetite.	<1	
352	Granite.			do.	do.	1/8	65		15			20		Trac	•	Trace							Biotite. Magnetite.	Trace.	
352	Tonalite and granite.	None.		branching.	do.	1/8-1/4	55		30			15		Trace		Trace.							do.	<1	
353	do.	do.	Crosscutting	Lenticular	. do.	1/4	55		25			20		<1		Trace							Biotite.	Trace.	
354	do.	do.	do.	branching.	do.	1/8	70		15			15		Trace		Trace							do.	Trace.	
355	do.	do.	do.	do.	do.	1/2-3/4	58		25			15		2		Trace.									
356	do.		Not exposed.	do.	do.	1/4	65		15			20		Trace		Trace				63			Biotite. Magnetite. Beryl. 3		.1/8-1
357	Tonalite.		do.	Irregular.	Wall some.	1/64	85		<1			15													,
358	Hornblende gneis		do.	Lenticular	. One unit.	1/4	60		20			20		Trace		Trace							Beryl. 10	crystale	.1/2-
359	do.		do.	do.	do.	1/4	60		20			20		<1		Trace.							do. 4 Columbite- tantalite.	crystals	
360	Coarse-grained granite.	Hone.		Irregular.	do.	3/4	45		40			15		Trace		Trace.								Trace.	
361	Tonalite.		Not exposed.	Lenticular branching		1/2-3/4	45		30			25		<1		Trace.									
362	do.		de.	Lenticular	. do.	1/4	65		15			20		<1		Trace.									
363	do.		do.	Lenticular branching		1/2	55		25			20		<1		Trace							Beryl. 4	rystals.	1/2
364	do.		do.	Lenticular	do.	1/4-1/2	50		25			25		<1		Trace.							do. 10	rystals.	1/8-1
365	do.		do.	do.	do.	1/4	60		20			20		<1		Trace.							do. 1	rystal.	1/4
366	do.		do.	Lenticular branching		1/8-1/4	72		8			20		<1		<1							40. 6	rystals.	1/8-1
367	do.		do.	do.	do.	1/4-1/2	53		25			20		5		Trace.					Die St		do. 1	crystal.	1/4
368	Tonalite.		do.	do.	do.	1/8	72		3			25		Trace		Trace.									
369	do.		do.	Lenticular	. do.	1/4	65	ATGS:	15			20		Trace		Trace					10-10				
370	do.		do.	do.	do.	1/8-1/4	69		10			20		1		Trace.					Sales!				

Ausber and	Wall r	ock													Pog	matite	ye suffi						le adh			
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture								E	92.9		M	neralo	er Ta							
(P1. II)			Wall Fook		• tracture		Plagio	clase	Pert	hite	Grap		Que	rts	Muso	ovite	Gar	net	Tours	aline	Li thium	miners	ı.	Other	r mineral	•
							Percent	Size	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Size	Per-	Sigo	Per-	Sigo	Mineral	Per-	Sise	Kineral	Per- cent	(1.834
371	Tonalite.		Not exposed.	Lenticular	. One unit	1/8-1/4	64		15				20		1		<1									
372	do.		40.	do.	do.	1/8-1/4	67		12				50		1		Trace				1			1		
373	do.		do.	do.	do.	1/4-1/2	54		25				20		1		<1									
374	do.		do.	do.	do.	1/8	70		5				25		Trace											
375	do.		do.	Lenticular	- do.	1/8-1/4	70		10				20		<1		Trace									
376	do.		do.	Lenticular	. do.	1/4	50		30				20		<1		Trace									
377	do		do.	40.	40.	1/8-1/4	70		10				20		<1		Trace							Beryl. 1	crystal.	1/4
378	do.		do.	do.	do.	1/4	59		20				20		<1		1			de dele			0.2.5	do. 9	crystal	. 1/4
379	Tonalite and fine-grained granite.		do.	do.	do.	1/8	73		5				22		<1		<1							do. 1	orystal.	1/16
380	Tonalite.		do.	Lenticular branching		1/4	68		12				20				Trace									
381	do.		do.	Lenticular	. do.	1/4	65		15				20		Trace		Trace									=\=
382	40.		do.	Lenticular branching		1/8-1/4	70		10				20		<1		<1							Beryl. 4	orystal	. 3/4
383	do.		do.	do.	do.	1/8-1/4	65		15				20		<1		<1							Martite.	Trace.	
384	do.		do.	do.	do.	1/4-1/2	60		20				20		<1		<1							Beryl. 1	crystal.	1/2
385	do.		do.	Lenticular	. do.	6-8	7		53		33		40		Trace									Biotite. Columbite tantalite	Trace.	ale.
386	do.		do.	do.	do.	1/8	72		8				20		<1		Trace									
387	Coarse-grained granite.	None.		do.	do.	1/4	60		25				15				<1									
388	do.	do.	100 ° 4 10 10 10 10 10 10 10 10 10 10 10 10 10	do.	do.	1/8	75		10				15				<1									
389	Tonalite.	do.	Crosscutting	Lenticular	- do.	1/8	73		7				20		Trace											
390	Tonalite and granite.	do.		Lenticular	. do.	1/8-1/4	70		10				20		<1		<1							Beryl. 1	erystal.	1/4
391	Tonalite.	do.	Crosscutting	Lenticular branching	-Vall sone	<1/64 4-8	75 15				~		25 85		Trace				Trace							
392	Covered.		Not exposed.	do.	One unit.	1-2	10		40				20		<1		Trace.									
393	do.		do.	Oval.	do.	1	40		35				25		<1											
394	do.		do.	Irregular.	do.	1/4	63		7				25		5		Trace									
395	do.	100 M	do.	Lenticular	. do.	1/8-1/4	57		15				25	3 11	3	V.	-							18 6 7 7		

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J	7	6	
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	Wall:	rock													Pog	matite										
Aumber and name of pag- matite	Type and formation	Alteration	Relation to	Shape		Texture								A de la			M	neralo	ey .							
(Pl. II)	IOTMA-CION		wall rock		structure	inches		oclase	Pert	thite	Gray	phie nite	Qua	rts	Muso	ovite	Gar	net	Tours	aline	Li thius	miner	rje	Othe	r mineral	•
		4			7		Per- cent	Size	Per-	Size	Per-	Size (inches)	Per-	Size	Per-	Size (inch	Per-	Size	Per-	Size	Mineral	Per-	Size (inches)	Mineral	Per- cent	Sis-
396	Covered.		Not exposed.	Lenticular branching	One unit	1-2	149		40		1		10		1											
397	do.	13.	do.	Lenticular	. do.	1/8-1/4	70		5				25		Trace.		Trace					1				
396	do.		do.	do.	do.	1-2	38	100	45		20		15		2											
399	do.		do.	Irregular.	do.	1-2	50		30		1		20		41		Trace							Bietite.	(1	
400	do.		do.	Lenticular	. 40.	3/4	50		30				20		<1		Trace									
401	do.		do.	Irregular.	do.	1/8-1/4	70		10				20											Biotite.		
402	do.		do.	Lenticular	. 40.	3/4	140		45				15		Trace.		Trace									
403	do.		do.	do.	do.	3/4	50		35				15		<1		Trace.									
707	do.		do.	do.	do.	3/4	45		30				25		<1		Trace									
405	Tonalite.	None.		Lenticular branching		3/4	55		15				30													
406	do.		do.	Lenticular	. do.	1/2	47		30	6			20		3											
407	Hornblende gneis	•	do.	do.	do.	3/4	50	THE S	20	4			30		<1								7			
406	do.		do.	Lenticular branching		1/2	60		10				30		<1											
409	do.		do.	Lenticular	do.	3/8	70						30		<1											
410	do.		do.	do.	Vall some. Gore.	1/8	85		15	4			15 85		<1											
411	Quartzite.		do.	do.	Vall sone. Core.	6/4	85 30		39	1-10			15		1 2											
412	Hornblende gneis	. .	do.	do.	One unit.	3	53	1/8	35	10			10		2									Beryl.	0.3	
413	do.		do.	40.	Vall some. Core.	1/8	72 15		12				15		8 <1									Beryl.	€0.05	
414	do.		do.	do.	One unit.	1/2	##		20				35		1											
415	do.	Barrier .	do.	do.	do.	1	57		8	3			35				<1	3								
416	do.		do.	do.	de	3/8	55	51.	35				10		<1		Trace.							Beryl. 4	prystals	- 3/4
417	do.		do.	Irregular.	Wall some. Core.	1/8	55		15	4-6			35		5									do. 1	prystal.	3/4
418	do.		do.	Lenticular	One unit.	1/4	61		15				20		4									do. 1	crystal.	1/4
419	do.		do.	Lenticular- branching.	do.	1/2	75		15				10		<1		Trace.									
420	do.		do.	Lenticular	do.	3/8	75		10				15				Trace.							Beryl. 2	crystals	1/4
H21 Ben-Kauf No.	4 do.		do.	do.	do.	1/2	44		20	No.			35		1								M. F. G.		1	

Aumber and	Wall r												12.45	9	reg	mat1te				600			Suggist.			
name of peg- natite (Pl. II)	Type and formation	Alteration	Relation to wall rock	Shape	Internal structure	Texture inches)											M:	ineral	OEA			HA DA				175%
(P1.11)							Plagic	clase	Pert	hite	Grap		Qua	rtz	Musc	ovite	Gas	rnet	Tour	naline	Lithium:	mineral		Other	mineral	•
							Per- cent		Per-			Size inches)								Size	Mineral	Per-	ize nches)	Mineral	Per- cent	(1nche
422 Ben-Kauf Ho	l Tonalite.		Not exposed.	Lenticular branching	.wall	3	22		35				35		8											
					Central unit.	1/2	181		15				2								Lepidolite.	. 2		Microlite.		1. 1/4
					Footwall layer.	1/4	50		5				740		8											
423	Hornblende gneis	•.	do.	Lenticular	. One unit.	1/2	49		35				15		1											
424 Bazooka	Tonalite.		do.	do.	Wall sone.	1/2	1 ₅₂ 1 ₁₅ 1 ₂₀						10								Lepidolite.	. 8				
					Core.	1/2	120						85 60								Lepidolite, Spodumene. Amblygonite	12		Microlite.	1 cryst	al.
425	Hornblende gneis	•	do.	do.	One unit.	1/4	60		15				25											Beryl. 3	crystals	. 3/4
426	do.		do.	do.	do.	3/8	54	1/8	10	4			35		1											N. S.
427	do.		do.	do.	do.	1	30		40				30													
428	do.		do.	do.	do.	1	30		40				30		41					N. S.				Beryl. 2	crystals	. 1/2
429	do.		do.	do.	do.	1/2	70		15	8			15		<1											
430 431	Hornblende gneie		do.	Oval.	One unit.	1/2	55	<1/8	20	2			25		Trace.											
431	ao.		۵۵.	Lenticular branching	. wall layer.	1/4-1/2	30 65		20				25		<1									Beryl.	<0 .05	1
432	do.		do.	do.	Border sone Hanging wal	. 1/8	92		5 30				3 20		4		Trace							Beryl.	0.1	1/4-3/ 1/4-3/
					layer. Footwall layer.	1/4	62	1/8	100				35		1										crystal.	-, - ,,
433	do.	None.	Conformable.	Lenticular	.Hanging wall	1/2	种		20	4			35		1		Trace							do.	<0.05	1-1/
					layer. Footwall layer.	1/8	64						35		1											
434	do.	A - 6 ()	Not exposed.	do.	One unit.	1/4	54		8				35		3	1/4								Beryl. 1	crystal.	6
435	do.		do.		Hanging wal layer. Footwall layer.	1 1	40 65		35				15 35		2 <1									do. 4	crystale	
436	do.		do.	do.	Wall sone. Gore.	1/4	65 5	1/8	5 80	3			30 15		<1 <1									Beryl. 1	crystal.	1
437	do.		do.	do.	One unit.	1/4	46	1/8	20	8			32		2									do. 1	crystal.	1/2
438	do.		do.	do.	do.	1/4	55	1/8	15	6			30		<1									do. 1	crystal.	3/4

Number and	Wall r	ock				E-0/188	1								.Peg	matite		100/570		2006						
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure(Texture					Service II						N	ineralo	ey							
(P1. II)	101111111111111111111111111111111111111		Wall Tour		atticoure(Inches,	Plagi	oclase	Pert	hite	Gray	hic ite	Que	rts	Musc	ovite	Gaz	rnet	Tours	naline	Lithium	miner	d.	Other	mineral	•
							Percent	Size (inche	Per-	Size	Percent	Size (inches	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Size inches	Mineral	Per-		Mineral	Per- cent	Size
439	Hornblende gneis		Not exposed.	Lenticular	One unit.	1/4	55		15				30		<1									Beryl. 2	crystal	1/4
1410	do.		do.	Oval.	Hanging wal layer. Footwall	1 3	25 52		40 15	12			25 26		10 7									do. 2 Apalite.	Trace.	
441	do.		do.	Irregular.	Vall sone.	1/4	65		30	1			5 90		Trace										9	
7115	do.		do.	Lenticular		1/8	45		15				35		5											
1413	do.		do.	do.	do.	1/8	38		35				25		2											
Jijiji	do.		do.	do.	do.	1/4	50	1/8	15				35													
445	do.		do.	do.	do.	8	35	1/8	5				60	1												
446	do.		do.	do.	do.	1/4	53	1/8	7	2			40	1	<1		Trace.									
448	do.		do.	do.	do.	1/4	58		35				25		71											
14149	do.		do.	Irregular.	do.	1	40	1/8					57	2	2		Trace.			1						
450	do.		do.	Lenticular	. do.	1	34	1/4					65	1	1											
451	do.		do.	do.	do.	1	40		40	4	5-10		20													
452 Brown Derby No. 1	do.	None.	Crosscutting.	branching		4-6	145						52		3		<1		<1		Curved lepidolite.	<1				
					Vest branch Core. Albite- monarite- columbite pegmatite		86	1/8-1/					5	1/8			<1	1/4-1/	2 <1	1/1-1/	2			Gahnite. Biotite. Monasite. Golumbite- tantalite. Fluorite. Betafite.	1.4 Trace.	1/: 1/: 1/4-1- 1/4-21
			i.		Hanging wal layer. Perthite- albite- luarts pag- matite.	1 12	30		140				20		10				Trace					Beryl.	0.1	1/2 by
					langing wal Quarts pod	1. 24	125						70								Lepidolite.	5	3-4			
					curved lepi	1- 4-5	144						40		<1				<1	1/2-3	Curved lepi	a- 15	1/4-2	Topas.	1	4-8
					epidolite- microlite	1	143						15								Lepidelite.	140	1/32-1/	8 Microlite Topas. Beryl. 1	. 0.35	1/64-1

unber and	Wall r														Pe	gmatite										
ame of peg-	Type and formation	Alteration	Relation to wall rock	Shape	Internal structure	Texture inches)						-					M	neral	ey .							
(P1.II)							Plagi	oclase	Pert	hite	Graj graj		Qua	rts	Muse	covite	Gar	net	Tours	aline	Lithium	minera	n.	Other	mineral	•
-							Percent	Size inches	Per-)cent (Size	Per-	Size (inches	Per-				Per-				Mineral	Per-	Size (inches) Mineral	Per- cent	Size (inches
h52 Brown Derby No. 1(Con tinued)	Hornblende gneis	s. None.	Crosscutting	branching		4-6	125	1/2-3					55	1-10	1	1/2-1			<1		Lepidolite	. 10	1-7	Topas.	10	Up to across by 42 long. 1-4
					Footwall layer. Albite peg- matite.	1/4	90						8		Trace		<1		2					Biotite.	Trace.	
453 Frown Derby No. 2	do.	do.	do.	Lenticular	.Wall sone.	1/4-1/2	62 1 ₅₃						30 40		7		<1 <1		1		Lepidolite	. 5		Bictite. Topas. do. Microlite.	1 <1 1.1	4-g
454 Frown Derby No. 3	do.	do.	do.	branching	Forth end South end wall sone.	1/2-3/4	10 37		74 2	12-24	75		15 55	1-2	1 4	1-3	<1		<1					Biotite. Beryl. 2	2	
					South end	5-6	147		1				38				<1		3		Lepidolite	7			<1 Trace.	tsl.
455 rown Derby lke Ho. 4	Hornblende gneis and biotite schiet.	do.	do.		Southern and norther	1/8	82		10				8		<1		< 1									,
					Southern part core.	24-36 24-36	10		8 0 5				10 95		frace.									Beryl. 3	crystal	. 3-12
456 rown Derby lke No. 5	Hornblende gneis		Not exposed.	do.	fall sone. Core.	1/2	156		20				15		<1 1		<1							Microlite.	1 cryst	1. 1/4
457 rown Derby lke No. 6	do.		do.	Lenticular branching	Core.	1/4	72 164		5				20 30		5		<1 <1		<1		Curved lepidolite. Curved lepidolite.	3		Beryl. 1	crystal	1/4
455 own Derby ke No. 7	do.		do.	Lenticular.	Hanging wall layer cotwall layer.	6 1/4	10		70 10	Up to			20		4		<1							Beryl. 1	crystal	. 2
usg own Derby ke No. 8	Hornblende gneise and biotite schi	None.	Crosscutting.	Lenticular branching.	Wall some	3	167		35	4-12	2		20		3		<1		Trace		Lepidolite.	2		Gahnite. Beryl. 4 Topas.	frace. crystals	1/4-7
0wn Derby ke No. 9	Biotite schiet.	do.	do.	Lenticular.		2	45		##	Up to			10		1											
461 own Derby ke No. 10	Biotite schist and hornblende gneiss.	do.	do.	Lenticular- branching.	40.	3	PTF		35	Up to			20		1		Trace		Trace					Beryl. 8	crystal	. 1/2-
462	fornblende gneiss	. do.	Conformable.	Lenticular.	Wall some Core.	1/8-1/1	171		5				20		2 <1		2 <1	1/16	<1 <1	1/4-1/2				Biotite. Lepidolite	<1	1/16-3

Number and	Wall r	ock							-	BULL TO		1200		Pe	gmatite	•			1000			Marine 1		- 3	
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture										M	ineral	DEX.							
(Pl. II)	Tornacion		wall rock		structure		Plagiocla	se Pe	erthite		aphic anite	Qu	arts	Musc	covite	Gas	rnet	Tours	naline	Lithiu	n miners	10	Other	mineral	•
							Per- Si cent (inc	Per hee)cer	Siz	e Per	- Size	Per-	Size	Per-	Size	Per-	Size inche	Per-	Size	Mineral		Size inches	Mineral	Per- cent	Sise (inches
463	Hornblende gneis		Not exposed.	Lenticular	Hanging wall layer. Footwall layer.	3 1/4	25 71	52				20		j 1		<1 <1							Beryl. 1	erystal.	1/4
464	do.		do.	do.	Wall sone.	1/4-1/2	145	10				40 25		5									Biotite.	<1	
465	do.		do.	do.	Vall zone.	1/4	70	15	Up	to		15 75		<1									do.	Trace.	
466	do.		do.	do.	One unit.	1/2	65	20				15									TOTAL VIEW				
467	do.	None.		do.	Hanging	4	30	40				30		<1		Trace.									
					Footwall layer.	1/4	65	15				20									0 NE 1				
					filling.							100													
468	Coarse-grained granite.	do.		do. I	Core.	3/4-1 3-4	3	77 60		30		20		<1											
469	do.	do.		do.	ne unit.	3/4-1	3	72				25		Trace		Trace.							Biotite.	Trace.	
470	do.			do.	do.	1-2	5	75				20		<1											
471	do.			Lenticular- branching.	do.	1/4-1/2	25	49		10-	15	25		1											
472	do.	None		do.	do.	2-3	4	69		30		25		5		Trace.									
473	do.			Lenticular	do.	3/4	30	7171		25		25		1		Trace.									
474	do.	None.		Lenticular branching		3/4	3	77		30		20		<1		Trace.				17 A22 18					
475	do.	do.		Lenticular,	do.	1-2	1	78		20		20		1											
476	do.	do.		Lenticular- branching.	do.	1-1/2	2	78		30		50		Trace		Trace.							Biotite.	Trace.	
477	do.	do.		do.	do.	1/4-1/2	50	25				25		<1		Trace.							Beryl. 21 Magnetite.		.1/8-1/
478	do.	do.		Lenticular.	do.	3-4	1	73		90		25		1											10
479	do.			Lenticular- branching.	do.	1	2	77		50		20		1		Trace.									
450	do.			Lenticular.	do.	1-2	5	75		70		20				Trace		10.10					Magnetite.	Trace.	
481	do.			do.	do.	1-2	2	78		60		20		<1		Trace									
482	do.		E A	do.	do.	1-2	1	79		75		20		<1											
483	do.			do.	do.	1/2-3/4	15	65		30		50		Trace									Beryl. 1 Biotite.	crystal Trace.	1/4
484	do.	None.	TANK S	do.	Wall zone. Core.	1-2	2	78 30		90	4	20 70	1/8										Magnetite.	Trace.	

	Wall z	rock									100			Pe	gmat1te	•									
Mumber and name of peg-	Type and	Alteration	Relation to	Shape	Internal											M	neralo	ey.			55			441	
matite (Pl. II)	formation		wall rock		structure	inches)	Plagic	clase	Perti	hite	Grapi	hic i to	quarts	Mus	covite	Gar	net	Tours	aline	Li thiur	miners	ale .	Other	mineral	
							Per-	Size	Per-	Size (inche	Per-	Size Pe	- Siz	Percent	Size (inche	Per-	Sise	Per-	Size (inches	Mineral	Per-	Size	Mineral	Per- cent	Size (inches
485	Coarse-grained granite.	Hone.		Lenticular	. Main uni Fracture filling.	6	5		75 V		85		0										Biotite.	Trace.	
486	do.	do.		Lenticular	- Main uni . Fractur filling.	. 3-4	5 Trace.		75 5		85	4 5	5	Trace		Trace.	1						Biotite.	Trace.	
487	do.			Oval.	One unit.	3-4	10		70		70		,												
488	do.			Lenticular	. Main uni Fracture filling.	. 3-4	15		70		80	,	5	Trace		Trace.		Trace.							
489	do.	None.		Lenticular branching	One unit	4	3		82		90	,	5			Trace.							Magnetite.	Trace.	
490	do.			Lenticular	. do.	2-3	3		77		90	2	,		i se								Magnetite.	Trace.	
491	do.			do.	do.	3	2		78		95	2				Trace.							do.	Trace.	
492	do.			Lenticular	- Main unit	. 3-4 5-6	1		79 35		95	8	5										do.	41	
493	do.	None.		do.	One unit.	4-5	Trace.		80		95	2	,			Trace.							Nagnetite.	Trace.	
494	do.	do.		do.	do.	4	3		77		95	2				Trace.							Biotite. Magnetite.	Trace.	
495	do.			do.	do.	1	5		75		80	2				41							Biotite.	Trace.	
496	do.	None.		do.	do.	3-4	15		55		1	. 3													
497	do.	do.		Irregular.	do.	3	10		65		1	2				Trace.									
498	do.	None.		do.	Main unit. Fracture filling.	1-2	35		35		1	3													
499	do.			Lenticular	One unit.	5	23		47		1	3		Trace.											
500	do.			Lenticular branching	do.	2	30		45		1	5	1	Trace											
501	do.			Lenticular	do.	5	8		50		1	3		2		100						et alle		NAC PROPER	TOWN ST
502	do.			do.	Wall some.	1/4	10		60	1,12	1	3				Trace.			Sec.						
503	do.			Lenticular branching	One unit.	6	7		65	Up to		10		3	3/4										
504	do.			Oval.	fain unit. Fracture filling.	2	30		140	Up to		10		5											
505	do.			Irregular.	Main unit. Fracture filling.	2 1	30		45	4	140	10	5			-									

humber and	Wall r	ock		100	A SULMAN				100		100	323			Pog	matite							1000			
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture											MI	neral	ogy							
(P1. II)	TOTALLION		Wall Fook		acructure (inches	Plagic	clase	Pert	hite	Gran	phic nite	Qua	rts	Musc	ovite	Gar	net	Tour	naline	Li thium	miners	10	Other	minera	10
							Per-	Size inches	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Sise	Per-	Size	Mineral		Size inches)	Kineral	Per-	Size (inche
506	Coarse-grained granite.	None.		Lenticular	- One unit.	4	15		55		85		30		Trace.		Trace.									
507	do.	do.		Oval.	Main unit. Practure filling.	2	20		55		80		25 100		<1											
508	do.	do.		Irregular.	Main unit. Fracture filling.	1-2	20		60				20		Trace.											
509	do.	do.		do.	Main unit. Fracture filling.	1/2	40		30		1		30 100		<1		Trace.									
510	do.		Not exposed.	Lenticular	. Main unit Fracture filling.	. 1/4	70		5				25													
511	do.	Fone.		do.	Main unit. Fracture filling.	1/2-1	30		40 10		1		30 90													
512	40.	do.		Irregular.	One unit.	1/2	30		39		50		30		1											
513	do.		Not exposed.	do.	do.	1/4	33		35		50	1/2	30		2											
514	Hornblende gneiss and granite.		do.	Lentiuclar	. do.	4	15		50		85		35													
515	Hornblende gneis		do.	Irregular.	do.	4	15		50		85		35													
516	Hornblende gneiss and granite.	•	do.	Lenticular	do.	2	34		35		-		30		1											
517	Hornblende gneis		do.	Irregular.	Main unit Fracture filling.	2-3	45		25				30 100		<1											
518	do.		do.	do.	One unit.	1-2	20		50		40		30													
519	do.		do.	Lenticular	do.	1/4	55		15				30											Motite.	Trace.	
520	do.		do.	Irregular.	Wall sone. Core.	1/4	50		20		1		30 90		Trace		Trace.									
521	do.		do.	Lenticular	One unit.	1/4	65		5				30													
522	do.		do.	do.	do.	3/8	80						20													
523	do.		do.	Lenticular- branching.		2	50		30	Up to	1		20													
524	do.		do.	Lenticular.	Wall zone.	3/4	10		65		~		25													
525	do.		do.	do.	One unit.	1/4	57		25				18										10 Sept			
526	Dacite.	None.		do.	do.	1/4	55		20	Up to			25		4 1					in him					A COLUMN	100

hunber and name of peg-	Wall ro	Alteration	Relation to	Shape	Internal	Texture									Pe	gmatite	8/10/10	neral	Dev .							
(Pl. II)	formation		wall rock		structure (inches)	Plagi	oclase	Pert	hite	Grap	hic	Qua	rts	Mus	covite		rnet		naline	Lithium			Other	mineral	
							Per-	Size	Per-	Size	Per-	Size	Per-	Size inches	Per-	Size	Per-	Size (inche	Per-	Size	Mineral)	Per-	Size (inches)	Mineral	Per-	Size (inches)
527	Dacite.	None.		Lenticular	. One unit.	1-1/2	20		50				30					46.0								
528	Hornblende gneis		Not exposed.		do.	1/4	65		25	3/4			10				<1									
529	do.		do.	Lenticular	. do.	4		184.6	1		85		~													
530	Hornblende gneis and granite.		do.		- Main unit	. 1-2	20		45				35 100													
531	Coarse-grained granite.		do.	Irregular.	One unit.	1	25		40				35													
532	do.		do.	Lenticular	Fracture	. 1/4	50		25	Up to			25 100				Trace									
533	Dacite.	None.		do.	filling.	1/4	53		25	Up to			20		2	3/4						de de				
,,,					Core.	,	"		-57	4			100		•	3/4										
534	Hornblende gneis	. do.		Irregular.	Wall zone. Core. Fracture filling.	1/8-1/4 4-12 6	58 7 7		15 62 62				25 30 30		2 1 1									Beryl. 1	crystal	. 4
535 Brown Derby No. 5	Tonalite.	do.		do.	fall sone. Intermediat sone.	1/8-1/4	161 35		10				2 5 55		1		Trace.	Up to 1-1/	<1 <1 2		Lepidolite do.	Trace 5		Beryl. Apatite. Columbite- tantalite. Microlite.	Trace.	by 3/4
					, Core.	6	20		39				40		1						Lepidolite	Trace		Topaz.	<1 <0.1 2 crysta	1/4 4-6 1/2-3-1 1s. 1/1 1-1/2
536	Hornblende gneis	. do.		do.	Wall sone.	1/4-1/2	45		10	Up to			40		5											
				Res :	Core. Fracture filling.	24-36	10	1/4	55 25				25 75		5									Biotite.	5	8
537	do.		Not exposed.	do.	Wall zone. Core.	1/4 8-10	58 10		65				35 20		2 5									Beryl.	0.3	1/4-3
538	do.		do.	do.	Wall zone.	1/4	57		10	Up to			25		8	1/4										.10. 1
539	do.		do.	Lenticular.		6 1/8	8 69		80				50 12 20		3	2/4								Beryl.	0.5	1/64-6
540	do.		do.	do.	One unit.	1/4	62	1/8	15	Up to		WIE	20	1/8	3	1/4	1000	1				1				

Number and	Wall r	ock	THE STREET						15.61						Pegmat	lte										
name of peg-	Type and	Alteration	Relation to	Shape	Internal structure	Texture											Mine	eralo	gy							
(Pl. II)	formation		Wall rock		structure	inches	Plagi	oclase	Pert	hite	Graph		Quarts		Muscovi i	te	Garne	et	Tours	aline	Lithiu	miner	als	Othe	r mineral	•
							Percent	Size	Per-	Size inches	Per-	Size Pinches)c	er- Si	lze P	er- Sin	e Po	er- S	Size	Per-	Size inches	Mineral	Per-	Size (inches)	Mineral	Per- cent	Size (inche
541	Hornblende gneis	s. None.		Lenticular	. Wall sone	. 1/4	60 5		10 25				25 70		5		<1							Beryl.	Several	3/4
542	do.		Not exposed.	Irregular.	Wall zone.	61/4	60 53		30 5				10		<1 <1										pieces.	
543	do.		do.	do.	Wall zone. Core.	1/2	55		10			1	30		5									Beryl. 1	crystal.	1-1,
544	do.		do.	Lenticular	One unit.	1/8	68		10				20		2											
545	40.		do.	do.	Wall zone.	1/8-1/4	69		10				20		1 2		<1									
546	do.		do.	do.	Wall zone. Core.	1/4	55 5		25				20		<1 <1									Biotite.	Trace.	
547	do.		do.	do.	One unit.	6	25		30				43		2											
548	do.	None.	Crosscutting.			10	15		50			N M	33		2									Beryl. 3	prystale	3-4
				branching	Footwall	1/8	74		3				20		3											
					layer. Consolidate	1/2	63		15				20		2											
549	do.			Lenticular	. Wall zone Core.	. 1/4	58 10		15 20				25		2											
550	do.		Not exposed.	Irregular.	One unit.	1/4	53	1/8	50				25		2									Monasite.	crysta	4. 3/1
551	do.		do.	Lenticular	. do.	1/2	54		25	Up to			20		1									Beryl.	crystal	. 1/2
552	do.		do.	do.	do.	1/8	60		10				30		<1										0.56	
553	do.		do.	do.	Wall zone.	1/8	70		15				14		1											
554	do.		do.	do.	One unit.	1/2-3/4	50	1/8	25	Up to			25		<1											
555	do.		do.	do.	Hanging wal layer. Footwall layer.	1/8	25 70		60 15				15	•	1	Tr	ace.									
556	do.			Irregular.	Main unit. Fracture filling.	1/4	70 20		15 55	Up to		TO THE PERSON	15													
557	do.			Lenticular	Wall zone	8	19		20	Up to			30		1											
558	do.		Not exposed.	Lenticular branching	Fracture				20	Up to			20		1	Tr	ace.									
559	do.		do.	Lenticular	filling. Main unit	6-12	5 65	1/8		Up to		9 2 3	30		4 1/	/4										

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36.1	
F	100
V41	1982
page 2	2
	1000

Mumber and	Wall ro	ck													Peg	atite										
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture											M	neralo	ey							
(P1. II)	TOTALCTOR		Wall Fock		structure	licite,	Plagic	clase	Pert	hite	Gray	hic ite	Quar	tz	Musoc	vite	Gaz	rnet	Tour	maline	Lithiu	m miner	als.	Othe	r mineral	le .
							Per-					Size									Mineral	Per-	Size	Mineral	Per-	Size
560	Hornblende gneis	s None.		Irregular	Wall some Main core. Small pod.	4	60 20 5	1/8	4 50 75				36 30 20		<1 Trace.	+1	Trace							Beryl.	0.45	1-1/2-
561	Hornblende gneie	•		Lenticula	. North en South end.	1. 1/4	60		20				20 25		Trace		Trace									
562	Coarse-grained granite.		Not exposed.	Irregular	Wall sone Core.	1/8	93		1 35	Up to			6 45		<1 <1		Trace									
563	Hornblende gneie			Lenticular	. One unit	3/4	45		40	Up to			15													
564	do.		Not exposed.	Irregular	Main unit Fracture filling. do.	1/4	60		20 50	Up to	~		19 50 100		1		Trace									
565	do.		do.	Lenticula: branching	- Wall zon	. 1/8	70 35		5 35				25		Trace.		Trace									
566	do.		do.	Irregular	Main unit Fracture filling.	1/4	50		20		~		30 100													
567	do.	Hone.		do.	Wall zone.	1/5	78		2				20		Trace.											
568	do.	do.		do.	Main unit. Fracture filling.	1/8	67		7 50	1/2-4			25		<1											
569	do.		Not exposed.	Oval.	One unit.	1/8	50		20				30		Trace.			-								
570	do.	None.		Lenticular	. do.	1/8	67		8	Up to			25		<1		Trace								4	
571	do.		Not exposed.	Irregular	do.	1/8-1/4	73		5	6			20		1		1									
572	do.		do.	do.	Wall sone. Core.	1/4	65		10 9				25 90		<1		<1									
573	do.	None.		do.	Main unit. Fracture filling.	1/4	64		15 50				20 50		1		Trace									
574	40.	do.		do.	Main unit. Fracture filling. do.	1/8-1/4	54		15 50				30 50 100		1		<1									
575	do.		Not exposed.	Lenticula	Core.	1/8	60		15				25 100		<1											
576	do.		do.	Oval.	Wall some	1/8	73 20		2 45				25 35		Trace.	V.	<1							V.		
577	do.		do.	do.	Main unit. Fracture filling. do.	1/5-1/4	66		5 50				25 50		3		1									

Aunber and	Wall r	ock					N.C.								Pog	matite						7000				
ame of peg-	Type and	Alteration	Relation to	Shape		Texture										1.772	M	neralo	EY							
(Pl. II)	formation		wall rock		structure	(inches)	Plagio	clase	Pert	hite	Gray	hic ite	Qua	rts	Musc	ovite	Gar	net	Tours	aline	Lithium	miner	de .	Other	mineral	
							Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size Inches	Per-	Size inches)	Mineral	Per- cent	Size inches)	Mineral	Per-	Size (inche
578	Hornblende gnei		Not exposed.	Lenticular	. One unit	1/8-1/4	61		1				35		3		Trace									
579	do.		do.	Irregular.	Main unit. Fracture filling.	1/8	75		7				15 100		3		Trace									
580	do.		40.	Oval.	One unit.	1/8	80		15				5		<1	100	Trace									
581	do.		do.	Irregular.	do.	1/8-1/4	63		8				27		2		Trace									
582	do.		do.	Oval.	Wall sone.	1/8	70		10				20		Trace.		<1									
583	do.		do.	Lenticular	. One unit.	1/2	40	1/8	30	4-6			30		<1		Trace									
584	do.		do.	Irregular.	do.	3/4	37		25				35		3		Trace									
585	do.		do.	Lenticular	. do.	1/2	30		39	Up to			30		1											
586	do.		do.	do.	do.	3/4	25		45				30		<1			-								
587	do.		do.	Irregular.	do.	1	49	1/8	20	2-6			30	2-6	1		Trace.									OG-E
588	do.		do.	Lenticular	. do.	1/4	53		20	Up to		7	25		2											
589	do.		do.	do.	do.	1/4	32		35	1-2			30		3							2.70		Beryl.	Several	
590 eryl and	do.		40.		Wall sone.	1/4	55		20				25		<1										crystale	. 1/4
Rare Mineral Lode					rone.	36	<1		50	12-60			20		30	Up to	<1					150		Beryl. Columbite- tantalite.	<0.05	1/2-8
					Core.	•							100											Gahnite. Unknown.	Trace.	<1/32 1/4
591	do.		do.	Oval.	One unit.	1/8	60		19				20		1											
592	do.		do.	Lenticular branching	- Wall sone	1/8	45		20 70	1/2			30 30	1/8	5	1/8-1/	2					199				
593	do.		do.	Lenticular	One unit	1/4-1/2	28		30				40		2											
594	Tonalite.	None.		do.	do.	1/2	56		18			Ju-Smelf	25		1		<1									
595	do.	do.		do.	Hanging wal	1 1-2	40		37				20		3		<1									
					layer. Footwall layer.	1/8-1/4	84		1				15		<1							in marke				
596	Hormblende gneis		Not exposed.	Irregular	One unit.	1/4	30		39	Up to			25		1											
597	do.		do.	Oval.	do.	1/2	25		39	Up to			35		1											
598	Quartzite.	None.	Crosscutting.	Lenticular	. do.	1/2-1	39		38				20		3		Trace						part and	Parmil 1		
599	Hornblende gneiss			Lenticular-	do.	1/4	44	1			PERM					*			1905					Beryl.	crysts	4. 3/

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humber and	Wall r	ock				- Unio			50000					Pe	mat1te									
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure			1974								Mineral	gy							
(P1. II)	Tormacton		wall rock		structure.		Plagio	clase	Pert	hite	Grap		Quartz	Musc	ovite	Garnet	Tour	naline	Lithium	minere	als	Other	mineral	8
								Size nches)	Per-	Size nches	Per- cent (Size inches	Per- Size	Per-	Size inches	Per- Size	Per-	Size inches)	Mineral	Per-	Size (inches)	Mineral	Per- cent	Size (inches
600	Hornblende gnei		Not exposed.	Oval.	One unit.	1/8	75		9				15	1										
601	do.		do.	do.	do.	1/16	70		14				15	1										
602	do.		do.	Lenticular	.Wall zone. Core pod.	1/8 1/2	64 30	1/16	20 50	2 3_4			15 20	1 <1	1/8		Trace	15				Beryl. 1	crystal.	. 1
603	do.		40.	Oval.	One unit.	1/32	79		5				15	1										
604 White Spar No. 1	do.	None.	Crosscutting	.Irregular.	Wall zone. Intermediat zone.	• 2	¹ 50		32 15	16			20	3	1-1/2				Lepidolite	. 5	2	Topas. Beryl.	<1 0.1	Up to 5
					Core. Small pod.	1/32			40	24			60						Lepidolite	90		Columbite- tantalite. Microlite.	Trace. Trace.	
605	Tonalite.		Not exposed.		Service of the service of	1/4	54		25	30			18	1	1/8			- 32	Lepidolite	155	1/)6	AICTOTICS.	0.1	
606	Hornblende gneis	s. None.	Crosscutting		do.	1/8	70	- North	15	3			15	<1	1/2									
607	Tonalite.		Not exposed.	branching	-Main unit. Fracture	1/32 1/8	75 50		9 23	3/4			15 25	1	1/16				Lepidolite.	. 2	1/8			
608	do.		do.	Lenticular		1/16	70		14	4			15	1	1/16									
609	Hornblende gneis		do.	Lenticular branching	- do.	1/32	70		9	3			20	1	1/16									
610	do.	None.	Crosscutting.	Lenticular	. do.	1/4	42		38	2			20	<1	1/32									
611	do.		Not exposed.	do.	do.	1/2	40		40	8			20	<1	1/8									
612	do.	None.	Crosscutting.	Lenticular branching	- Hanging wall layer Footwall	1/4	41 78		37 5	1			20	2	1/8									
613	do.		Not exposed.	Tantianlas	layer.	1/8				2			20											
					Footwall layer.	1/32	55		25				20	<1	1/16									
614	Biotite schist.	Mone.	Crosscutting.	do.	One unit.	1/8	65		19	3-1/	2		15	1	1/16									
615	Hornblende gneis	. do.	Conformable.	do.	do.	1/8	70		15	2-1/	2		15	<1	1/8									
616	do.		Not exposed.	Irregular.	Hanging	1/4	50		30	4			20	<1										1000
					wall layer Footwall layer.	1/32	75		10	5			15	<1	1/16	Trace.								
617	do.		do.	Lenticular	Hanging wall layer	1/2	30		60	5			10	<1	1/16									
	The same of the sa				Footwall layer.	1/16	64		20	4			15	1	1/16	Prace.								

Number and	Wall r	ock													Pe	matite						Mark W.				0.86 1900
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture		Huge									Mi	neral	ogy							
(P1. II)	Iormation		Wall rock		structure	Inches,	Plagio	clase	Pert	hite	Graz		Qua	rts	Muse	covite	Gar	net	Tour	maline	Lithium	minera	de .	Other	mineral	•
							Per-	Size (inches	Per-	Size	Per-	Size (inches	Per-	Size (inches	Per-	Size (inche	Per-	Size inche	Per-	Size	Mineral	Per-	Size inches)	Mineral	Per- cent	Sis (inch
618	Hornblende gneis	s. None.	Crosscutting		Hanging wall layer. Footwall layer.	1/2	54 74		25 10	5			2 0		1	1/16	Trace									
619	do.		Not exposed.	Lenticular		1/8	65		20	4			15													
620	do.		do.	do.	do.	1/8	70		15	1			15		<1	1/16										
621	do.		do.	do.	do.	1/4	60		25	3			15		<1	1/16					71-1-1					1
622	do.		do.	do.	do.	1/8	65		25	2			10		<1	1/16										
623 .	do.		do.	do.	do.	1/4	40		40	2			20		<1	1/8										
624	do.	None.		Irregular.	do.	1/16	60		15	2			54		1	1/32										
625	do.	do.		Lenticular	. do.	1/16	65	201	9	2-5			25		1	1/16	To the									
626	do.		Crosscutting	do.	do.	1/16	70		15	3-9	5.4		15		<1	1/16										100
627	do.		do.	40.	Wall sone. Core.	1/16	60 30		20 50	1-2 1-2			50		<1 <1	1/16										
628	do.		do.	do.	Hanging wall layer. Footwall layer.	1 1/32	70		35 10	1-2			20		1	1/4										
629	do.		Not exposed.	do.	One unit.	1/4	50		29	2-4		Hall I	20		1	1/8										
630	do.	None.	Crosscutting	do.	do.	1/4	64		20	3-6			15		1	1/16										
631	do.		Not exposed.	do.	do.	1/16	69		15	3-6			15		1	1/16	Trace			Starring .						
632	do.		do.	do.	40.	1/4	45		35	2-4			20		<1	1/16										
633	do.		do.	do.	do.	1/32	65		20				14		1		Trace									
634	do.		do.	do.	do.	1/64	65		20				13		2											
635	Tonalite.		do.	do.	do.	1/16	60		19				20		1											
636 White Spar No. 2	Hornblende gneis	. None.	Crosscutting.	do.	Wall zone. Core.	1/4 1/32	45		30 1 10	-1/2- 2	-1/2		20 35		5		<1				Lepidolite Lepidolite	<1	1/8 1/32	Fluorite. Microlite. Beryl. 2	<0.01 crystal	. 5/8
637	do.	do.		do.	Wall zone. Core pod.	1/4	120		38 35	8-12			20		2 5		Trace.				Lepidolite	Trace		Columbite-	a. Trac	
					Fracture filling.	1/4	¹ 50		5	4-10			30		15						Lepidolite	Trace		tantalite. Beryl. 1 Kicrolite.	crystal	1/8
638	do.		Not exposed.		Hanging wall layer Central	1/4	30		5				50		15											
					layer. Footwall	3	70						30													
				The second of th	layer.	1/8	64		5				30		1		1						S 107 20			

Number and	Wall r						TOTAL ST		-		ATTIO			-	Peg	matite							0.5			
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture									17. 6		M	neral	PEY	1000						
(Pl. II)							Plagio	clase	Pert	thite	Grap		Qua	rtz	Muso	ovite	Gar	net	Tour	naline	Lithium	minera	1.	Other	mineral	
		4					Per- cent	Size		Size	Per-	Size inches	Per-	Size inches	Per-	Size im he	Per-	Size (inche	Per-	Size (inches	Mineral	Per-	Size inches)	Mineral	Per- cent	Size (inche
639	Hornblende gneis	•.	Not exposed.	Lenticular	.Wall sone. Core pod.	1/4	59 10		20.	3-4 8-15			2 0 67		0.5	1/2			W							
640	do.		do.	do.	One unit.	1/8	50		30	4-7			19		1	1/2										
641	do.		do.	Irregular.	do.	1/16	65		14	3-6			20		1	1/8										
642	do.	None.		do.	do.	1/16	55		23				20		2	-/-	Trace									
643	Hornblende gneis and biotite schist.	do.	Crosscutting.	do.	Wall sone. Core pod.	1/8 1/4 1/4	72 20 19		7 55 60				20 15 20		1 4		Trace.							Beryl. 2	crystals	. 1/2
644	Hornblende gneis		Not exposed.	Lenticular			43			10-12					1											
645	do.	None.	Crosscutting.		Wall sone Core. Small pods	. 3/16	55 19		24	1-1-1	/2		20 40		1 1	2	O.5									
646	do.		Not exposed.	Lenticular			49		20 50	2-3 1-2 Up to			30 40		1 2	3-4										
647	do.		do.	do.	Wall sone. Core pod.	1/64	75		13 80	10 4-7 9-14			10		2 5	1/4			Trace							
648	do.		do.	Irregular.	One unit.	1/16	75		8	5-7			15	2			Trace.									
649	Hornblende gneise mica schist, and quartzite.		do.	Lenticular	Wall sone. Core pod.	1/32	54 30		10 47	6-8			35 20		1 3		Trace.									
650	Hornblende gneiss		do.	Irregular.	One unit.	1/8	69		10	4-6			20		1											
651	do.		do.	do.	do.	1/8	69		10	2-1/2			100			1/8										
652	do.		do.	Lenticular.		1/8	40	1. S. T	20	12-14	,		20		1	1/8										
					Footwall layer.	1/16	70		9	2-3			50		1	1/8	Trace									
653	do.	None.		Lenticular- branching.	One unit.	1/16	60		20	5-9			20		0.5	Up to	frace.									
654	do.		Not exposed.	Oval.	do.	1/16	60		10	3-5-1	2		29		1	1/16	Trace.		. 170							
655	do.	None.		Irregular.	Wall sone.	1/4	45		35	8-12			19		1	Jp to			100				A			
					Core.	2	10		52				35		3	2	0.5									Up to
656	do.			Lenticular	. One unit.	1/16	60		18	2-4			20		2	1/8							154			
657	do.	None.		Irregular.	Wall zone. Core pod.		31 65		40 15	4-10			25		4	3/8										
658	do.	do.		Lenticular	. One unit.	1/16	50		19	3-4			30		1	Play										
659 Suckhorn and Feldspar lodes.	Hornblende gneis and tonalite.	do.	Croascutting	Irregular	. Wall sone Intermediat	.1/4-1/2	59		20 50	1-3			20		1 40	4-12	Trace							Biotite. Beryl Columbite- tantalite. Monasite.	Trace. Several	crysta

humber and	Wall r						The same				-				regi	matite										
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure(Texture inches)					2				No.			neralo								
(Pl.II)							Plagi	oclase	Pert	hite	Grap	hic	Quart	E	Musoc	ovite	Garı	net	Tours	aline	Li thium	minera	ls	Other	mineral	
							Percent	Size inch es	Per-)cent(Size inches	Per-)cent(Size Pe	nt(ir	Size nches)	Per- cent(i	Sise	Per-	Size	Per-	Size (inche	Mineral		Size inches)	Mineral	Per- cent	Size
659 Buckhorn and Feldspar lodes (Continued)	Hornblende gneis and tonalite.	None.	Crosscutting.		. Core, Buckhorn claim. Core pods, on ridge.	2 ¹ 4	1		91 53				7		1 1	1								Beryl.	0.7	1/2-4
660	Hornblende gneis		Not exposed.	Lenticular	Wall sone. Core pods.	1/16	70		9	3-6			20		1	1/8										
661	do.			Irregular.	Wall zone. Intermediat zone. Core.	1/8 • 1/2	70 20		10 149	5-#		1	8 0	W. W.	2 10	1/8								Beryl.	0.5	1/2-2-1
662	do.		do.	Oval.	One unit.	1/8	78		4	Up to			5		3		<1									
663	do.		do.	Irregular	Wall some	. 1/4	53		20 65				5		2											
664	do.		do.	do.	Wall some	1/4	70	1/4-1/2		Up to			0		2 <1	,	frace.									
665	do.		do.	Oval.	One unit.	1/8	87			Up to		E. 22		p to	<1		1									
666	do.		do.	do.	Wall sone. Core.	1/8	76 25		2 35			2 4	5		2		<1									
667	do.		do.	Lenticular	Wall sone	1/8	80		5 10			1	5		Frace.		<1									
668	do.		do.	Lenticular- branching.		1/8-1/4 2-3	72		8				0		<1 <1	2	race.									
669	do.		do.	Irregular.	One unit.	1/4-1/2	.60		10			3	0		<1											
670	do.	None.		do.	do.	1/4	60		15	1-4		2	5		<1	-	race.									
671	do.		Not exposed.	do.	Main unit. Fracture filling. do.	1/2	56		19 50			5 10			1		race.									
672	do.		do.	Lenticular	. One unit	1/8	71		8		7	2	0		1		Trace.					4.29	NAME OF			
673	do.		do.	do.	do.	1/4	40		25	1000		3	5		<1		1.14			- 10						
674	do.			Irregular.	Wall sone.	1/8	65 15		10 45			2 3	5		<1 5	Up to	<1							Columbita- tantalite	Trace.	
675	do.		Not exposed.	Oval.	Main unit. Fracture filling.	1/8-1/4	70		15 50	Up to		1 5			<1		Trace.									
676	do.			Irregular.	Main unit.	1/8	70		10			2			<1		<1									
					filling.	4	3	197				9	5	1	2	1	1			17/19			100	S TRANS		1 1000

umber and	Wall r	ock													Peg	matite				100	Jan					7 -19
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure(Texture								0-0			M	neral	ogy .							
(P1. II)	I OF ME CION		ANTI LOCK		structure		Plagio	clase	Pert	hite	Grap	hic	Qua	rts	Musc	ovite	Gar	net	Tour	naline	Lithium	minere	als	Other	nineral	Le .
							Per-					Size inches)								Size (inches	Mineral	Per- cent(Size inches)	Mineral	Per- cent	Sise (inche
677	Hornblende gneis	•.	Not exposed.	Oval.	Main unit.	1/8-1/4	60		15	Up to			25		<1									Bioti te.	Trace.	
					Fracture filling.	1							100													
678	do.		do.	Irregular.	Main unit. Fracture	1-1/2	40	1/4	35	3-4			25	3-4	<1		Trace.									
					filling.	6-8	15		55				100 30									gial)				
679	do.			Lenticular branching	- Wall sone	. 1/8 4-8	72 15		12				15		1 <1		Trace		He							
680	do.		Crosscutting.	Irregular.	Main unit.	1/8	74		5				20		1		<1									
					Fracture filling.	1							100													
681	do.			do.	One unit.	1/16	70		9				50		1		Trace.								15	
682	do.		Crosscutting.	Irregular	Core pod.	3/4	70 15		14 54	3-5 6-10			15 30		1	1/4	Trace.									
683	do.		Not exposed.	Lenticula	r. One unit	. 1/16	70		9	4-5			20		1	1/32										
684	do.	None.	Conformable.	Irregular	Hanging wall layer Footwall layer.	1 1/32	45 75		38				15 15		2		Trace.									
685	do.			Lenticular		1/2	30 65		19	4-6 2-3			20		2	1/2	Trace.									
686	do.		Not exposed.	Oval.	One unit.	1/4	60		20	5-7			19		1	1-1/2	Trace.									
567	do.		do.	Lenticular	do.	1/4	50		39			W	10		1											
588	do.		do.	Irregular	do.	1/32	70		14	2-4			15		1	3/4										
689	do.		do.	Oval.	Hanging wall layer Footwall layer.	1/2	30 45		55 33	6-12 2-3			15 20		<1 2	1/2										
590	do.		do.	do.	One unit.	1/32	60		23	5-7			15		2	1/4		317								
i91	do.		do.	do.	do.	1/4	65		20	3-4			13		2	1/2							2000			
92	do.		do.	do.	do.	1/4	40	1	40				19		1											
93	do.				Wall zone. Core.	1/16	60 20		24 19				15 60.		1											
94	do.			Irregular.	One unit.	1/64	74		10		~		15		1											
95	do.		Not exposed.	Lenticular	. One unit	1/4	50		30				20													
696	do.		do.	do.	do.	1/8	60		20			19	20													
697	do.		do.	Irregular.	do.	1/8	30		54	4			15		1	Ca A			100							

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	,100mg	
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	C	2

fumber and	Wall r	ock										1355	8 1		Pog	mati te		9			The state of the s					H. THE
ame of peg-	Type and	Alteration	Relation to	Shape	Internal structure	Texture					Trips						Mi	neralo	EY							
(Pl. II)	formation		wall rook		structure	Inches	Plagic	clase	Pert	hite	Grap	hic ite	Qua	rts	Musc	ovite	Gar	net	Tours	aline	Li thius	miner	le	Other	mineral	
							Per-	Size inches	Per-	Size nches	Per-	Size	Per-	Size inches	Per-	Size	Per-	Si se inches	Per-	Size inches) Mineral	Per-	Size	Mineral	Per-	Sise (inches
698	Hornblende gnei		Not exposed.	Irregular.	One unit.	1/16	60		19.5				20		0.5											
699	do.		do.	do.	do.	1/8	50		30		30	8	20		Trace		Trace	1/16						Biotite.	Trace.	1/8
700	do.		do.	do.	do.	1/4	55		30		30	5	15											Magnetite.	Trace.	1/1
701	do.		do.	do.	do.	1	20		50		1	10	30		1 2 3		Trace	1/16						do.	Trace.	1/4
702	do.		do.	Lenticular	. do.	2	35		50		140	10	15				Trace	1/32						Magnetite.	Trace.	1/10
703	do.		do.	Oval.	do.	1/4	50		30		1		20	Te	1-7		Trace							Biotite.	Trace.	
704	do.		do.	Irregular.	do.	1/8	50		35		30		15										-			
705	do.		do.	Lenticular	. do.	1/16	65		15				20											Biotite.	Trace.	
706	do.		do.	do.	do.	1/8	140		30				30										10000			R
707	do.		do.	Oval.	do.	1/16	70		5				25													
708	do.		do.	Lenticular	. do.	2	10		65		80		25					-1.5								
709	do.		do.	do.	do.	5	10		65		80		25													
710	do.		do.	do.	do.	1	30		140				30												135.5	
711	do.		do.	do.	do.	3	10		70		80		50													
712	do.		do.	Oval.	do.	3	10		60		1		30													
713	do.		do.	Irregular.	do.	3	10		60				30													
724	do.		do.	Lenticular	. do.	3	10		60		80		30													
715	do.		do.	Irregular.	Wall zone. Intermediat	2	10		60				30													
					zone. Core.								50 100		50	3/4			0							
716	do.		do.	Lenticular	One unit.	3	10		70				20				Trace.									
717	do.		do.	do.	do.	2	10		70				50		9140	1.30										
718	do.		do.	do.	do.	3	10		70		80		20					in a				150	- 3			
719	do.		do.	Oval.	do.	1/2	60		10	8			30	130			Trace.	1 74								
720	do.		do.	Lenticular	do.	1/4	70		10				20													
721	do.		do.	Oval.	do.	1	10		60				30				Trace.									
722	do.		do.	Lenticular	do.	1/8	70		10				20													100
723	Hornblende gneis and coarse-grain granite.	d	do.	do.	do.	1/4	50		30				20				Trace.							Biotite.	Trace.	
724	Hornblende gneis		do.	do.	do.	1/4	60		10				30													
725	do.		do.	do.	do.	1/4	50		20				30			-5-93	No.									1

Aumber and	Wall r	ock													Peg	matite										
ame of pag-	Type and formation	Alteration	Relation to	Shape	Internal structure(Texture											Min	eralo	gy							Title!
atite (Pl. II)	IOFMETION		wall rock		structure		Plagic	clase	Perth	ite	Graph	hio ite	Quart	is .	Musc	ovite	Garn	et	Tours	aline	Lithium	minera	de .	Other	mineral	
							Per-	Size Inches	Per-	Size nches	Per- cent(Size nches)	Per-	Size ches)	Per-	Size	Per-	Size	Per-	Size inches	Mineral	Per- cent	Size inches)	Mineral	Per-	Sise (inches
726	Hornblende gmei	10.	Not exposed.	Lenticular	. One unit	1/4	50		30			-57	20													
727	do.		do.	do.	do.	1/4	60		10				30	-												
728	do.		do.		- Wall mone .Intermedia sone. Core.		10 167		65		75		25 30 20		3	2								Beryl. 1	crystal.	3/8
729	do.		do.	Lenticular	.Wall sone. Intermediat sone. Core.	1/4	70		10				80 60 80													
730	do.		do.	do.	One unit.	1/16	60		10				30											Biotite.	Trace.	
731	do.		do.	do.	do.	1/8	30		50		1		20													
732	do.		do.	Oval.	40.	1/32	50		15				5													
733	do.		do.	Lenticular	. do.	5	20		60		60		20													
734	do.		do.	do.	do.	1/16	60		10				30													
735	do.		do.	Lenticular branching	- do.	8	30		70		50		50													
736	do.		do.	do.	do.	1/4	70		10				20													
737	do.		do.	Lenticular	. do.	1	10		70		50		20													
738	do.		do.	do.	do.	2	10		70		60		20													
739	do.		do.	Lenticular branching	- Wall sone	1/4	60		20 80				50													
740	do.		do.		Wall sone. Intermediat sone. Core. Southwest dike.	1/4	30		50 70				20 30 100 20				Trace.									
741	Hornblende gneis and granite.	. None		Lenticular	. One unit.	1/32	70						30											Biotite.	Trace.	
742	Hornblende gneis		Not exposed.	Oval.	do.	1	10		60		~		30				Trace.									
743	do.		do.	Lenticular	. do.	2	10		70		60		20			2014		1								1000
744	do.		do.	Lenticular branching	- Wall sone .Intermedia sone. Core.	1/16 1/16 1/4	70 35 30		10				60		5	2										
745	do.		do.	Oval.	One unit.	2	10		70		70		20		18.5			1								
746	Granite.		do.	Lenticular		2	10		70		70	ATTENDED IN	20													

Munber and	Wall r	rock						0.23		Phy	1	ELEKS!			Pe	gmatite	•			No. 1	Barrier,	Mari			10.020	7243
name of pag- matite	Type and formation	Alteration	Relation to	Shape	Internal	Texture											N	ineral	- KY							
(Pl. II)	IOFMACION		WALL FOOK		structure	inches)	Plagi	oclase	Pert	hite	Grap		Qua	rts	Muse	covite	Ga	rnet	Tour	maline	Lithiu	m miner	als	Other	mineral	
							Percent	Size	Per-	Size	Per-	Size inches	Per-	Size	Per-	Size (inche	Per-	Size	Per-	Size inches	Mineral		Size (inches)	Mineral	Per- cent	Size (inache
747	Hornblende gneis		Not exposed.	Lenticular	. Wall sone Intermediat sone. Core.	. 1/8	30 20		50				20 70 100		10	1-1/4										
748	do.		do.	do.	Wall zone. Intermediat zone. Core.	1/8 3/8	70 50		10				20 45 100		5	1-1/2										
749	do.		do.	do.	One unit.	2	10		70		60		20													
750	do.		do.	do.	Wall sone. Intermedia sone. Core.	1/8 te 3/8	70 50		2	5			26 45 100		5	1-1/3										
751	do.	400	do.	do.	One unit.	1/8	30		50				20						1000							
752	do.		do.	do.	Vall sone. Intermedia sone. Core.	2 3/8	10		70		70		20 75 100		15	1										
753	do.		do.	do.	One unit.	1/8	60		10				30				Trac									
754	do.	Soft Control	do,	do.	do.	1/8	40		20	N. O.			30	80.49			Trac									
755	do.		do.	do.	do.	1/16	50		10		1		40													
756	do.		do.	do.	do.	1/16	65		15	5			50													
757	do.		do.	do.	do.	1/8	60		15				25													
758	40.		do.	do.	do.	3/8	10	100	60				30													
759	de.		do.	do.	do.	1/32	60		10	- 53			30				Trace									
760	do.		do.	do.	do.	1/32	60		25				15				Trace									
761	do.		do.	do.	do.	1/4	70		10				20				98			100						
762	do.		do.	do.	do.	1/8	60		10				30							14.8						
763	do.		do.	do.	do.	1/4	70		10				20				Trace	•								100
764	do		do.	do.	do.	1/32	70		5				25				Trace	1 37/3				1				1
765	do.		do.	do.	do.	1/32	70		5		15,000		25				Trace	-								
766	do.		do.	Oval.	do.	1/4	70		10				20	No.												100
767	do.		do.	Lenticula		1/4	50		50		60		30													
768	do.		do.	Oval.	do.	3/8	10		70	19/16			20										1			
769	do.		do.	do.	40.	1/8	35	1	35				30					1000					3/ 10	Biotite.	Trace.	
770	do.		do.	Lenticula		1/8	30		40		50		30	- 10						1			The same	do.	<1	
771	do.	1	do.	do.	do.	1/4	30	130	50		50		20	19 19	1	173			- 12		BERRY				-	13 34

	133
2/4	
1/4 1/2	
1/32	
2	0
2 3	C
. 2-1/2	
	1000

Monber and	Wall re	ook													Pog	matite	•									
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture											361	neral	EY							
(P1. II)	IOFMETICE		WELL FOOK		structure		Plagic	clase	Pert	hite	Gray	hic	Que	rts	Muso	ovite	Gar	net	Tours	maline	Lithium	minera	10	Other	mineral	
							Percent	Sise	Per-	Size	Per-	Size	Per-	Size inches	Per-	Size	Per-	Sise	Per-	Sige inches) Mineral	Per-	Size inches)	Mineral	Per-	Sise (inche
772	Hornblende gnei	٥.	Not exposed.	Owal.	One unit.	1/4	140		20		10		20													
773	do.		do.	Lenticular	. do.	1/8	50		20		15		30											Biotite.	Trace.	
77%	Coarse-grained granite.	None.		Lenticular		. 3 4	25		49.5	12	50	8	15 50											Magnetite Biotite. do.	Trace. Trace. 0.5	1/4
775	do.	do.		Lenticular	. One unit	1	50		30	3	10	5	80				Trace	. 1/16						do. Nartite.	Trace.	1/1
776	do.	do.		do.	do.	1	10	No.	69.5		1		80				Trace							Biotite.	0.5	1/3
777	do.	40.	7.4	do.	One unit.	1	50		30		NO		80				Trace							Biotite. Magnetite.	Trace.	
776	Coarse-grained granite and	do.		Irregular.	Hanging wall layer.	18	10		70		70	24	29											Biotite.	1	
	hornblende gneis	•.			Footwall layer.	5	70	1/2	10	1	5		17		Trace.	1/8	Trace	. 2						do.	3	2
					Core.	12			60			-	140												Trace. Trace.	3
779	Hornblende gneis		Not exposed.	Lenticular	. One unit.	5	50		15		35	8	35				1	,								
780	do.		do.	do.	Hanging	5	10		60	3	70	8	30											Biotite.	Trace.	
					Vall layer. Footwall layer.	1	70		5		10	3	25											Martite. Biotite.	<1 Trace.	
781	do.		do.	do.	Wall sone. Core.	1-1/2	5		65 10	3	75	5	30 89		Trace.	2								Martite.	<1	
782	de.		do.	Irregular,	Wall some. Core.	3/4	4.5		60	3	75	5	35 80		1	1	Trace.							Martite.	0.5	
783	Coarse-grained granite and hornblende gneis	Mone.	Crosscutting.	do.	Core.	1 3	30		50 65		60		30		5		Trace							Biotite. Martite. Beryl. 2	Trace.	. 2-1
784	Hornblende gneis		Not exposed.	Lenticular	. One unit.	1	25		55		70		20				Trace							Martite.	Trace.	
785	do.		do.	do.	do.	1/16	140		35	3	140		25	10			Trace.									
786	do.		do.	Lenticular branching	- do.	1	10		70		60		80													
787	do.		do.	Lenticular	do.	2	10		65		80		25				Trace.									
788	do.		do.	do.	do.	3	10		75				15													
789	do.		do.	do.	do.	3	10		75				15													
790	do.		do.	Lenticular branching		3	10		65		80		25		Prace.		Trace.				4					
791	do.		do.	Oval.	do.	2	10		65		80		25													
792	Granite.			Lenticular	do.	1/2	60		20		20	6	20		5139		P. S. S.		-			Sur!	BATE	Biotite.	Trace.	

Aumber and	Wall	rock	AND AND										280-010		Peg	matite										
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture									Top La		M	neralo	ey	Con Series						
(Pl.II)	107420101		Wall Fook		a cructure;	inches,	Plagi	oclase	Peri	thite	Grap	hic	Que	rts	Musc	ovite	Gaz	net	Tours	aline	Lithium	miners	18	Other	mineral	•
							Per-	Size	Per-	Size	Per-	Size	Per-	Size inches	Per-	Size	Per-	Size inches	Per-	Size inches	Mineral		Size inches)	Mineral	Per- cent	Size (inche
793	Granite.			Lenticular	One unit.	. 1	10		65		80		25													
794	do.			do.	do.	1/2	50		30		40	6	20									1.50		Biotite.	Trace.	
795	do.			do.	do.	1/2	35		45		40		20													
796	do.			do.	do.	1	20		55		60		25													
797	do.			Lenticular branching	One unit.	1	10		70		60	6	20													
798	40.			Lenticular	do.	2	10		70	33	60		20													
799	do.			do.	do.	2	10		60		50		30													
800	do.			do.	do.	2	10		60		50		30													
801	do.			do.	do.	1-1/2	15		60		50		25				Trace.									
802	do.			do.	do.	1/4	40		40		30		20													
803	do.			do.	do.	1/16	30		50		40		20													
804	do.	None.		Lenticular branching	do.	1/4	30		55		45		15													
805	do.	do.		Lenticular	do.	1/4	20		60		50		20				Trace.									
806	do.	do.		do.	do.	1	10		70		60		20											Biotite.	Trace.	
807	do.			do.	do.	2	10		65		55		25													
808	do.			Lenticular branching.	do.	1	10		65		5 5		25											Magnetite.	Trace.	
809	do.			do.	do.	2	10		75		65		15													
810	do.			Oval.	do.	1/2	15	S. C.	60		50		25													
811	do.			Lenticular	do.	1/4	10		60		50	CEL	30													
812	do.	None.		Lenticular- branching.	do.	1/2	50		55		45		25													
813	do.			Lenticular	do.	1/4	20		60		50		20											Biotite.	Trace.	
814	do.	None.		Lenticular- branching.	do.	1/4	15		65		55		20													
815	do.			Lenticular.	do.	1/16	35		40	1	1		25	1,211	357 July		Trace.									
816	do.			do.	do.	1/8	50		25		1		25													
817	do.			Irregular.	do.	1/8	30		45	7	10		25		951		Trace.									
818	do.			Oval.	do.	1/2	30		50		40		20				Trace.			841						PAGE 1
819	do.			Lenticular.	do.	1	10		70		60		20				Trace.							Biotite.	Trace.	
820	do.	100000		do.	do.	1	10	128	70		60	100	20	N. L.	186		Trace.		BIN W					do.	Trace.	-

per and		rock	TEN TO STORE							STATE OF THE PARTY OF					6	matite	100		100			AL STATE OF		Tells of the		Sales
of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure(Texture											M	neralo	ey							2
.11)							Plagic	oclase	Pert	hite	Grap		Qua	rts	Muso	ovite	Gar	net	Tours	aline	Lithium	miners	de	Other	mineral	
							Per-	Size	Per-	Size	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Size inches	Per-	Size inches	Mineral	Per-	Size inches)	Mineral	Per- cent	Sise (inch
521	Granite.			Lenticular	. One unit.	1/4	25		55				20													
255	do.			Oval.	do.	1/32	50		30				20													
323	do.			Irregular.	do.	1/4	45		35	1	25	7	20													
124	do.			Oval.	do.	2	10		60		50		30							- [-						
125	do.	Hone.		Lenticular	do.	1	10		65		55		25													100
126	do.	do.		do.	Wall sone. Core.	1/4	30 10	1/16	50 55	5 5	40	6 12	20		<1 15	1/2								Monasite.	Trace.	1/2
127	do.	do.		do.	One unit.	1/2	35	1/4	45	5	20	5	20		<1	1/8	Trace.							Monasite.	frace.	1/4
26	do.	do.		do.	do.	1/4	50		30		20	6	20							200.10				do.	Trace.	1/4
29	do.	do.		Irregular	do.	1/4	50		30	5	20	8	20				Trace.							Biotite.	Trace.	
30	do.	do.		Lenticular	do.	1/2	25		55		10		20	ly m												
31	do.	do.		do.	do.	3/4	35	1/16	45	3		8	20		Trace		Trace.						1			
32	do.	do.		Oval.	do.	3/4	10		65	2	40	6	25				Trace.							Biotite. Magnetite.	Trace.	
33	do.			do.	do.	1/16	60		24	5			15		1	2										
34	do.	None.		Lenticular branching.	do.	1/2	30		55		40		15		Trace.		Frace.							Biotite.	Trace.	
35	do.	do.		do.	do.	1/2	30		50	4	40	4	20				Frace.							do.	Trace.	
36	do.	do.		do.	do.	1/4	50		30		Gr.		50				Frace.									
37	do.			Irregular.	do.	1/4	20		65		55	16	15									10 Table				
38	do.	None.		Lenticular.	do.	1/8	45		40		30	8	15									111				100
39	do.	do.		Lenticular- branching.	do.	1/8	65		20		10	10	15													
40	do.			Lenticular.	do.	2	10		70		60	6	20													
41	do.			Oval.	do.	1/8	65		20	1			15		Trace									Martite.	Trace.	
42	do.			do.	do.	1/4	65		15				20		Trace											
43	do.			do.	do.	1/8	35		50		40		15											Biotite.	Trace.	
1111	do.		The same	Irregular.	do.	1/8	70		15	2			15													
45	Covered with glacial till.			do.	do.	2	30		1414				20		0.5	1	0.5	1/32								
46	do.			do.	do.	1	33-5		45		35	8	50		0.5	3/4	1	1/16								
	glacial till.																		2 1736							

	Wall r	ock													Pe	gmat1te										
Number and name of peg-	Type and	Alteration	Relation to	Shape		Texture											M	neral	gy							
(Pl.II)	formation		wall rock		structure (inches)		oclase	Pert	hite	Grapi	hic ite	Qua	rts	Muse	covite	Gas	rnet	Tours	naline	Lithium	minere	ıls	Other	mineral	
							Per-	Size inches)			Per-	Size									Mineral	Per-	Size (inche	Mineral	Per-	Size (inches
gu7 Black Wonder	Hornblende gneis coarse-grained granite, and quartsite.	s, Mone.	Crosscutting		Wall sone, south and west parts. Wall sone, north and east parts. Intermedia	1-1/2	24 50 79		60		50 less than 10	8	15 15		5	1/4-5	Trace	1/4-2						Martite. Biotite. Martite.	1 <1 Trace.	1/2-1-1/
					sone. Intermedia		55						15			Up to		1/						Monarite. Columbite-	17 cryst	als. 3/4
					Core. do. do.				1				100											tantalite.	Trace.	1/4
848	Hornblende gneis		Not exposed.	Lenticula	r. One unit	. 2	55		25	4,4			20											Magnetite.	Trace.	
849	do.		do.	do.	do.	1	49		30		20		20													7
850	do.		do.	Irregular.	do.	2	25		54		40		20				1									
851	do.		do.	Lenticular branching		1	35		45		35		20				<1									
852	do.		do.	Lenticular	do.	3/4	35		45	7	35	6	20				Trace							Biotite. Magnetite.	Trace.	
853	do.		do.	do.	do.	1	35		45	3	35	6	50				Trace.	1/32				6.27		Condition		
854	do.		do.	Oval.	do.	3/4	40		45	1	35		15				<1							Biotite.	Trace.	Him a
855	do.		Crosscutting.	Irregular	do.	3	15		65		55		20											Magnetite.	Trace.	
856	do.		Not exposed.	do.	Wall sone. Core.	3/4 5	10		70 50	6	60	5	20 50		Trace.									do.	Trace.	
857	do.		do.	Oval.	One unit.	3	50		60	3	55	6	20											Magnetite.	Trace.	1/16
	Hornblende gneis and granite.			Lenticular	do.	3	35	1/16	40	3	30	8	25											do.	Trace.	1/16
859	Hornblende gneis	•	Not exposed.	do.	do.	2	35		45	3	25	6	20				Trace.	1/2								
860	do.		do.	Irregular.	do.	1	30		50	8	45	3	50				Trace.							Biotite. Magnetite.	Trace.	
861	do.		do.	Oval.	do.	1/2	70		10		5		20							1						
862	do.		do.	Lenticular.	do.	1/4	60	1	20				20						1132							
863	do.		do.	do.	do.	1/16	70		10				20									1				
864	do.		do.	do.	do.	1/2	50		30		20		20													
865	do.		do.	do.	do.	1/2	50		50		40		30									186.8				
866	do.		do.	do.	40.	1/2	50		30		20		50			1						100	2			
867	do.		do.	do.	do.	1/4	60		20				20										S. T.	Magnetite.	Trace.	
868	do.		do.	do.	de.	1/8	60	18.0	10	1			30				Trace.				A STATE OF THE PARTY OF THE PAR	1				

4PO S FE0

amber and	Wall r												1100		reg	matite	8 734									
ame of pag-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture inches)				-			Aut				Mi	neralo	67							
(P1. II)							Plagio	clase	Perti	hite	Grapi	hie ite	Quar	rts	Muso	ovite	Gar	net	Tours	aline	Li thium	miner	de .	Other	mineral	ls
							Per- cent(Size Inches	Per- cent(Sise inches	Per-	Size	Per-	Size	Per-	Size	Per-	Sise inches	Per-) cent(Size Inches	Mineral	Per-	Sise inches)	Mineral	Per- cent	Sis
869	Hornblende gneis		Not exposed.	Lenticular	. One unit.	1/16	70		5				25				Trace									
870	do.		do.	Oval.	do.	1/16	60		20				20													
871	do.		do.	do.	do.	1/4	65		15		10		20				Trace									-
872	do.		do.	Lenticular branching	- do.	1/16	70		10		~		20													
873	do.		do.	Lenticular	. do.	1/4	55		20		10		25													
874	do.		do.	do.	do.	1	30		50		40		20				Trace	NA LI								
875	do.		do.	do.	do.	1/2	50		30		20		80				Trace			-1						
876	do.		do.	40.	do.	1/2	140		140		30		20				Trace		2					Biotite.	Trace.	
877	do.		do.	do.	do.	1/4	60		20		10		20				Trace									
878	do.		do.	do.	do.	1/8	50		30		20		20				Trace							Biotite.	Trace.	
879	Hornblende gneis	s Mone.	4.00	do.	do.	1/2	50		30		20		20				Trace									
880	Hornblende gneis	•.	Not exposed.	Leaticular	. One unit.	1/2	140		140		30		20													
881	do.		do.	do.	do.	1/8	60		10				30													
882	do.		40.	do.	do.	1/8	40		30				30									555		Biotite.	Trace.	
883	do.		do.	Oval.	do.	1/8	50		20		10		30													1.3
884	do.		do.	Lenticular	. do.	1/8	60		10	5			30											Biotite.	Trace.	
885	do.		do.	do.	do.	1/8	50		20		10		30		Trace											
886	do.		do.	do.	do.	1/8	50		20		10		30		Trace				No Today							
887	do.		do.	do.	do.	1/8	60		20		10		20													
888	do.		do.	do.	do.	1/8	50		20		10		30		Trace											
889	do.		do.	do.	do.	1/2	30		50		40		20													
890	do.		do.	do.	do.	1/8	50		20		10		30													
891	do.		do.	do.	do.	1/8	50		20		10		30							1500						100
892	do.		do.	do.	do.	1/2	30	Link !	50	100	40		20										Ben			
893	do.		do.	Oval.	do.	1/8	50	9	20		10		30													No.
894	do.		do.	Lenticular	. do.	3/8	140		28				30		2					1		19 18		1		1
895	do.		do.		Hanging wall layer. Footwall layer.	1/4	50 60		20				30 19		3											

unber and	Wall r	ook							1015	18 97				Pogmat	tite				de la constante					•	
ame of pag-	Type and formation	Alteration	Relation to	Shape	Internal	Texture										161	neral	ey.							
(Pl. 11)	Torantion		ANTI LOCK		structure		Plagioclas	e Pe	rthite		phie nite	Quar	ts N	uscovi	lte	Gaz	rnet	Tour	aline	Lithius	niner	als	Other	r mineral	le
							Per- Sissinche	Per	Size (Inche	Per-	Sise	Per-	Size Pe	r- Si	ise	Per-	Sise	Per-	Sise	Mineral	Per-	Size	Mineral	Per-	Sise (inche
896	Hernblende gnei	••.	Not exposed.	Lenticula	r. Hanging wall layer Footwall		55	8	5			20	Tr	ace.											
					layer.	1/32	75	!	1			20	22	ace.			5 1							1	
897	do.		do.	do.	Hanging wall layer Footwall	Maria 4	140	30				28	2												
					layer.	1/32	60	20	1 2 34			19	1								File	- A3			
898	do.		do.	do.	One unit.		35) ji				20	-												
899	do.		Crossoutting		do.	1	30	50				80			1/4			1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1							
900	40.		Not exposed.	do.	Hanging wall layer Footwall layer.	3/4	70	10				19.5	0.		/4					•					
901	do.		do.	40.	Hanging wall layer	1	20	64.5	C SO EN			15	0.												
					Pootwall layer.	1/4	50	lec	1/2			10	4	.1										-	
905	do.		do.	do.	One unit.	3/4	48	30	3			20	5	1	/4	Trace	•								
903	Rorablende gnei		do.	do.	do.	3/4	35	50	3			15	<	1											
904	do.	Hone.	Crosscutting	. do.	do.	1/2	30	50	3/4		133	20						File							
905	do.		Not exposed.	do.	Hanging wall layer Footwall	12002271	25	55				19.5	0.	1	/8										
					layer. Core.	1/4	50	30	14			20	Tre	ice.									Beryl.	Trace.	1/4-3/4
906	Hornblende gnet	•	do.	do.	Hanging wall layer Footwall	14	15	69	3			20	Tr	100. 1	/4	Trace									
					layer.	1/4	60	20	2			80	Tr	1	/16	Trace									
907	Hornblende gneis	s Hone.	do.	do.	One unit.	1/4	30	60	1/2			10													
908	Hornblende gnei		do.	do.	Wall sone. Intermedia	3/4	20	69	4	1		15		<1					1						
					core.	3/8	25	140	3/4			28.5	5	3,	/8	Trace							Beryl.	1.5	1/16-1
909	Hornblende gneis	s None.		do.	One unit.	3/4	20	50	1			30	Tre	ce. 1/	/4										
910	Coarse-grained granite.	do.		do.	Hanging wall layer. Footwall	3/4	5	80	2	10	3	15											Biotite. Nagnetite		
			25.5		layer.	1/4	55	35	1/2		1	10		41		Trace		15-11							
911	Hornblende gneis	•	Not exposed.	do.	Hanging wall layer. Footwall layer.	3/4	15 ks	69.5		3	3	15	0.	5											

Mumber and	Wall r							1				20.00			2.08	mat1te	San Park		-							
name of pag-	Type and formation	Alteration	Relation to	Shape	Internal structure										3,516		Mi	neralo	67							
(Pl. II)							Plagio	clase	Pert	hite	Grap	hie ite	Quan	rts	Muso	ovite	Gar	net	Tour	naline	Lithium	minera	ls	Othez	mineral	le .
							Per-	Size	Per-	Size	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Sise	Per-	Size inches	Mineral)	Per-	Size inches)	Mineral	Per-	Sise (inches
912	Hornblende gneis	8.	Not exposed.	Lenticula	. One unit.	3/4	25		60	2			15		<1		Trace									
913	do.		do.	do.	Hanging wall layer. Footwall layer.	3/4	15 75	1/64	65	2			20		Trace,		Trace.									
914	do.		do.	do.	One unit.	1/2	55		25	1			20	200	frace.	1/8							132.00		The same	
915	do.		do.	do.	do.	1/2	30		50	1-1/2			20		<1											
916	do.		do.	do.	Hanging wall layer. Footwall	1/4	30		39	3			30		1	1/2										
					layer.	1/64	60	on the	10				30				Trace.									
917	do.		do.	do.	Hanging wall layer. Footwall layer.	1/4	26 50		40				30		3									Biotite.	1	
918	do.		do.	do.	Hanging wall layer.	1/4	50		20				30		10	1	Trace.									
					Pootwall layer.	1/2	10		60	3			30		Prace.									Biotite.	Trace.	
919	do.		do.	do.	One unit.	1/4	28		40	2			30		2	1										
920	do.		do.	do.	Wall sone. Intermedia sone.	1/64 te 1/4	70		10				30		5	1	Trace.									
					Core.	1	5		59				34		5								1			
921	do.		do.	do.	Hanging wall layer Footwall layer.	1/4	60		70	3	-		20		2											
922	do.		do.	do.	Hanging	1/32	50		20				20		10	Mary 1	Trace.									
					wall layer Footwall layer.		70		20				10		10	1/2	Trace.									
923	do.		do.	do.	Hanging wall layer Footwall	1/4	48		20	4			20		2	1										
					layer.	1/4	55		14				30		1											
924	do.		do.	do.	Hanging wall layer Footwall		8		60	3			30		2											
000					layer.	1/32	19	1500	40				40		1					1						
925	do.		do.	do.	Hanging wall layer Intermedia sone.	1/8	30		50				60		25	1/2										
					Core. Footwall layer.	1 1/32	60		80	5			20				5. 3.									

Part Size					NAME.		Ser.	12	Buch	A 1986				0.0	10.0	0.00	200	-	100	230	100	644	1				33.2	egi.	43	Pe	gmat1t			3.5								
Page Canal Pertition Capable Pertition Capable Capab													xture	Textur	Pextur	exture	xture		NA.														Miner	ralo	gy							
Second S			hite	hite	hite	thite	Perthit	Pez	lase	oclas	agioc	Plag	nches	(inche	(lnche	inches	nches	Pla	lagio	loclas	ase	Per	erthi	lte				Qua	rts	Muse	covite		Garnet	t	Tours	naline	Lithiw	n miner	als	Othe	r mineral	le .
Second continues of the seco	ze	Per-)cent(i	Size	Size	Size	Size (inche	Per Si	Per	Size	Size (inc)	ent (Per-	t (inc	ize nche	Per	nt(in	Size aches	Per-	- Si t(inc	ize 1	Per-	Size	Per-	Size (inche	Pe:	r- Si nt(inc	ise	Per-	Size				Mineral	Per-	Size (inches
### granite. Polith 1/8				200			10	10			65	65	1/8	1/8	1/8	1/8	1/8	6	65			10	0					24		1	1/2	Tra	ce.							Monasite.	1 crys	tal. 1/2
928 Roreblende gestes. do. Lenticular. One unit. 1/4 ho 30 2 38 2 1/2																														<1										Biotite.	Trace.	
939 do.					1000	1000					40	40	1/4	1/4	1/4	1/4	1/4	140	40			1		2						2	1/2	134				NAME OF					-1600.	
930 40.			4	4	4	4	50	50			15	15	/4	1/4	1/4	1/4	14	1	15			50		4						5		1859										
1 1 1 1 1 1 1 1 1 1							14	14			60	60	1/64	1/64	1/64	1/64	1/64	60	60			14	•					25		1											1.5	
931 do. do. Lenticular Hanging branching vall layer. 932 do. do. Lenticular Hootwall 1/32 60 10 10 29.5 0.5 Frace. 933 do. do. Lenticular Hanging vall layer. 1/32 60 10 10 29.5 0.5 Frace. 934 do. do. Lenticular Hanging vall layer. 1/8 60 15 24 1 1 29 1 1 29 1 20 1/2 Frace. Protwall layer. 1/8 10 60 3 29 1 1 20 1 1 20 1 1 20 1 1 20 1 1 20 1							10	10														10)							10	1											
1 1 1 1 1 1 1 1 1 1							40	140			300 H		6.50	1/4			5.5					140					199			10	1	1100								Monazite.	1 cryst	11/4
Wall layer 1/8 60 15 24 1 1 1 1 1 1 1 1 1							10	10			60	60	1/32	1/32	1/32	1/32	/32	60	60			10					2	9.5		0.5		Trac	30.								1	
935 do. do. do. Hanging vall layer. Footwall layer. Footwall layer. 1/64 60 936 Fine-grained granite and horn-blende gneiss. 937 Granite. do. do. Lenticular. Hanging vall layer. Footwall lay																															1/2	Frac										
934 do. do. do. Hanging 1/4 30 40 4 28 2 1 935 do. do. do. Hanging wall layer. 7 Footwall layer. 1/64 60 10 28 2 936 Fine-grained grante and horn-blende gneiss. 937 Grante. do. Lenticular Ranging wall layer Footwall layer. 7 Footwall layer. 1/8 63 20 4 15 2 1/8 Frace. 1/16 Beryl Park Processing wall layer. 1/16 65 5 15 15 0.1 1/2																130						1					30	FIR		1		1										
934 do. do. do. Hanging wall layer Pootwall Pootwall layer Pootwall Pootwall layer Pootwall Pootwal			3	3	3	3	60	60														60		3						1		-										
935 do. do. do. Hanging 1/8 30 40 3 25 5 1-1/4 936 Fine-grained grante and horablende gneiss. 937 Grante. do. Lenticular One unit. 1/4 63 20 1/32 65 3 15 0.1 1/2 Lenticular Footwall 1 1/2 20 1/32 65 3 15 0.1 1/2			4	4	4	4	40 1	40						1/4								40		4						2	1											
			0		1	1	10	10			60	60	/64	1/64	1/64	1/64	/64	60	60			10						28		2												
936 Fine-grained granite and horn-bleade gneiss. 937 Granite. 40. Lenticular One unit. 1/4 63 20 4 15 2 1/4 Trace. 1/16 Beryl Dranching. 1-1/2 20 1/32 65 3 15 0.1 1/2			3	3	3	3	40	40														40		3						5	1-1/4	+										
granite and horn-blende gneiss. Dranching. Dranching						1.	_ ,	-					23.00	1000								_						Elle			- 0	1										
wall layer. Footwall								20			"	0)	/-	1/7	1/-	-/-		0)	0)			20						15		2	1/4	Tra	ice.	1/16						Beryl. 1	orystal	1/4
layer. 1/2 65 1/64 15 1 20 Frace.			3	3	3	3	55 3	65	1/32	1/3	80	20	-1/2			1-1/2	-1/2	20	20	1/3	/32	65		3				15		0.1	1/2											
			1	1	1	1	15 1	15	1/64	1/6	55	65	1/2	1/2	1/2	1/2	1/2	65	65	1/6	/64	15		1			1	50		Trace.											18/15	
938 Hornblende gneise. Not exposed. do. One unit. 1/2 25 65 2 2 5 10 <1		2	2	2	2	2	55 8	65			25	25	1/2	1/2	1/2	1/2	1/2	25	25			65	1	2	2	5	5	10		<1								1				
939 do. do. do. Hanging 3/4 40 35 2 10 4 20 2 1/2 Trace.									. 10					1							10																					
940 do. do. do. One unit. 1/2 25 60 2-1/2 14 1		5	Lo day	La de	La de	1000		3 30	1/64	1/6			9.46	1	7			18		1/6	/64	126		0 8 1	5	1	100			Trace	1/32	3				500		1				

Aumber and	Wall re	oak										23/10/	4. 9		Peg	gmatite		8-13								
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture											ж	neral	OET							
(P1. II)	Tormstion		WALL FOCK		structure	ladnes	Plagio	clase	Perti	hite	Grap		Qua	rts	Muse	ovite	Gaz	rnet	Tour	maline	Li thium	miner	ils	Othe	r minera	1.
								Size	Per-	Size inches	Per-	Size	Per-	Size	Per-	Size	Per-	Si se (inche	Per-	Size	Mineral	Per-	Size (inches)	Mineral	Per- cent	Sise (inches
941	Granite and hornblende gneise	None.		Lenticular	. One unit	1	15		55		5	5	30		Trace		Trace									
942	do.	do.		do.	do.	3/4	10		65	3			25													
943	Gneissic granite	. do.	Crossoutting	do.	do.	1/2	30	1/38	50	2			18		5	1/4										
944	Granite and horn blende gneiss.		Not exposed.	100 Jan 1	Hanging wall layer. Footwall	3/4	20		15	h.			15		15	1/8	Trace									
					layer.	1/4	50	1/32	10	5			20	1/4	20	1/4	0.1									
945	do.	None.			Hanging wall layer Footwall layer.	1-1/2	15	1/32	60	3			15		10	1/2										
946	Hornblende gneis	•	Not exposed.	branching		3/4	30		20	2-1/2			50		Trace,											
					layer.	1/2	50		140				10		<1		Trace.								-	
947	Granite.	Hone.			wall layer.	3/4	30	. /	50	2			20		Trace.											
948	Hornblende gneis		Not exposed.	do.	layer.	1/2	60	1/32		1			20													
,	normoreade gaste		not exposed.	40.	Hanging wall layer Footwall layer.		55		15				30		Prace.	,										
949	do.	ď	do.	do.	Hanging wall layer. Footwall	1/16	35		35				28		2											
					layer.	1/32	60		9				30		1											
950	do.		do.	do.	Hanging wall layer Footwall		40		30				28		2											
OE3					layer.	1/32	60		9				30		1		Trace.									
951	do.		do.	do.	Hanging wall layer Footwall layer.	1/8	70		140				17		3											
952	do.		do.	do.	Hanging	3/8	140		9 39	2000			50		1		Trace.			7. 7		24 14				
					wall layer Footwall layer.	1/32	60		19				20		1											
953	Hornblende gneis	•	do.	do.	Hanging wall layer	3/8	45		25				28		2											
					Footwall layer.	1/64	60		5				34		1		Trace.					55				
954	do.		do.	do.	Hanging wall layer Footwall	1/8	40		40				17		3								E NOT			
					layer.	1/64	70		9				20		i		Trace.					1			Tomas !	1

unber and	Wall r										1		-		reg	matite										Selfer I
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure(Texture inches)							7 5				Mt	neralo	er er						S Paris	
(P1. II)					- 1		Plagic	clase	Peri	thite	Graz		Qua	rts	Muso	ovite	Gaz	net	Tours	aline	Lithium	miners	10	Other	mineral	•
							Per-	Size inches	Per-	Size	Per-	Size Laches	Per-	Size inches	Per-	Size Inches	Per-	Sise inches	Per-	Size	Mineral	Per- cent	Size (inches)	Mineral	Per- cent	Sise (inches
955	Hornblende gneis	•	Not exposed.		wall layer		63		20		Ä		15		2											
956	do.		do.	do.	Hanging wall layer	1/32	70 58		20				20		2											
					Footwall layer.	1/32	74		9				15		1											To a chi
957	do.		do.	do.	Hanging wall layer	1/4	50		30				19		1											
					Pootwall layer.	1/32	70						30													
958	Hornblende gneis		do.	do.	Hanging wall layer Footwall	1/2	20		65	2-1/2			15													
					layer.	1/2	45		40	1			15				Trace.					A SE		Biotite.	Trace.	
959	Hornblende gneis		do.		Hanging wall layer Footwall		30	1/8	50	4			20		<1	1/4										
960	do.				layer.	1/4	70	1/32	10	1			80		Trace								7			
961	do.		do.		One unit.	3/4	45	- 1-6	40	2			15		<1						Feb. 1					
962	do.	Hene.	Crosscutting	do.	do.	1/2	55 30	1/16	20.00	3			19.5		0.5	1/4										
963	do.		Not exposed.		Hanging wall layer	1/4	48		50 20				30		2	1/2								Monagite.	1 cryst	1. 1/
		7.5			Pootwall layer.	1/32	70		4				25		1	e03										
964	do.		do.	do.	One unit.	1/16	50		20				30												5-7-11	
965	Coarse-grained granite and horn blende gneiss.	None.		Lenticular branching.	de.	1/8	50		20		~		30													
966	Hornblende gnei		Not exposed.	Lenticular	. do.	1/4	20		60		50		20													
967	do.		do.	do.	do.	1/4	30		50		40		20													
968	do.		do.	Oval.	do.	1/32	60		10				29				1									
969	do.		do.	Lenticular	. Hanging wall layer	1/4	10		60		5		25										7-2-12			in the same
					Footwall layer.	1/4	40		27				30	X 3	3											3
970	do.		do.	do.	One unit.	1/64	80			or in			20													
971	do.		do.	do.	Hanging	1/4	10		60				30					TEN					S. S.			
					Footwall layer.	1/64	60		5				30		5											

Sumber and	Wall r	ook													Peg	matite	•									
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture											M	ineral	EF							
(P1.II)	IOPMACION		wall rock		structure	(indnes	Plagic	oclase	Pert	hite	Grap	hic ite	Qua	rts	Musc	ovite	Gaz	rnet	Tours	naline	Lithius	n miners	de .	Other	mineral	•
							Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Sise	Per-	Size inches	Mineral		Size inches)	Mineral	Per-	Sise (inch
972	Hornblende gneis	•	Not exposed.	Lenticular	. Hanging wall layer. Footwall layer.	1/16	10		60				28		2											
973	do.		do.	do.	Wall some.	1/64	70 60		9				20 30		1 2											
974	do.		do.	do.	One unit.	1/32	60		13				25		2											
975	do.		do.	do.	Hanging wall layer. Footwall layer.	1/2	50		19				30		1	2										
976	do.		do.	do.	One unit.	1/64	75						25													
977	do.		do.	do.	do.	1/16	75						20		5											
978	do.		do.	do.	Hanging wall layer	1/64	10		60				30													
					Footwall layer.	1/16	50		20				22		3											
979	do.		do.	do.	Hanging wall layer Footwall	3/8	10		65				24		1											
					layer.	1/4	50		15				30		5											
980	do.		do.	do.	Hanging wall layer Footwall				70				28		2											
					layer.	1/64	70		5				25		3											
981	do.		do.	do.	Hanging wall layer Footwall layer.	1/8	10		60				30		h											
982	do.		do.	do.	Hanging wall layer	1/2	5		70				24		1											
					Footwall layer.	1/64	70	Envil I					28		2											
983	do.		do.	do.	Hanging wall layer Footwall	3/8	20		60				18		2											
					layer.	1/32	70		4				25		1											
984	do.		do.	Lenticular branching	- Hanging wall layer Pootwall	1/4	10		60				27		3											12.000
985	Hornblende gneis		do.	do.	layer.	1/4	30		10	-			45	Tan-E	15	7/1		. 10								
	and fine-grained granite.		40.		Hanging wall layer Footwall	•	20		50	5			50		10	3/4	Trace	. 1/64						Beryl. 1		1/2
					layer.	1/2	49.2		30	5			15		5	1/16	0.5	1/64						Monagite.	0.3 3 crysta	
986	do.		do.	Lenticular.	One unit.	2	15		50	12	1	5	19.5		15	3/4	0.5	1/32				112				

	Wall r	oak			DELT'S				1381	7			200		Peg	mat1te	•		Skell						-	
Number and name of pag-	Type and	Alteration	Relation to	Shape	Internal	Texture	100										M	ineral	KN						5.4.4.4	
(F1. II)	formation		wall rock		structure	inches)	Plagi	oclase	Pert	hite	Gran		Quar	ts	Muso	ovite	Gas	rnet	Tours	aline	Lithiu	n miner	als	Othe	r mineral	•
							Per-	Size	Per-	Size	Per-	Size	Per-	Size inches	Per-	Size	Per-	Size	Per-	Size	Mineral	Per-	Size (inches)	Mineral	Per-	Size
987	Hernblende gneis and fine-grained granite.		Not exposed.	Lenticular branching.	- Hanging wall layer Footwall layer.	3/4	45		10	4 2			15		10	1/4	Trace	1/16								
988	do.		do.	do.	Hanging wall layer Footwall	2	20		50	5			20		10	3/4	Trace	1/64								
					layer.	1/16	60		15	1		1000	15		10	1/16	Trace	1/64		DE O						
989	do.		do.	do.	Wall sone, Core,	1/64	75 40		10 29	1 4			15		Trace	. 1/8	<1	1/128						Beryl. 2	prystal	1/4-3
990	do.		do.	Lenticular	. One unit.	1/8	45		35	2			20		Trace	. 1/4			500						1000	
991	Hornblende gneis		do.	do.	do.	3/8	54		30	4			15		1	1/16	<1	1/126								
	Hornblende gneis and fine-grained granite.		Not exposed	Lenticular branching.	vall layer. Footwall	1/4	40		30	2			20		10	1/4										
					layer.	1/32	73		10	1			15		5		Trace.									
993	do.	do.	do.	Lenticular		1/8	64		50	1-1/2			15		1		Trace.	PIE C								
	Hornblende gneis		do.	do.	do.	1/4	65		20	3			15		Trace		<1	1/128								
995	do.		do.	do.	do.	3/8	45		40	3			15		Trace									Beryl. 3	crystals	1/4-1
996 997	do.		do.	do.	do. Hanging wall layer	3/4	35		45	5	5	3	15 25	EGG	Trace	1/2	<1	1/128								
			- 1 a a a		Footwall layer.	1/4	75		15				10		<1									Monasite.	1 orvata	1/2
	Fine-grained grau and hornblende gneiss.	ite None.		do.	One unit.	1/8	55		25	2			50		Trace	.1/32								Beryl. 4	3 3 3 3 6	1000
999	Hornblende gneise		Not exposed.	Lenticular branching	do.	1/8-1/4	55		25				50		<1		Trace.							do. 1	crystal.	1/4
1000	Fine-grained granite and horn- blende gneiss.	None.		Lenticular.	Wall sone Core.	1/5	63		20 50	4	10	5	15 20		2 10	1/16 3/8	<1 Trace.	1/128								
1001	do.		Not exposed.	do.	One unit.	1/2	35		45	3			20		Trace	1/4	Frace.	1/64						Beryl. 1	erystal.	1
1002	Hornblende gneiss		do.	do.	Wall sone.	1/64	78		5				15		Trace	1/8	2	1/64								
			Market 1		Core.	1/2	29		50				20		1	1/4	frace.	1/32						Beryl. 3 Columbite tantalite	-	100
1003	blende gneiss.	None.	-		Hanging wall layer Footwall		20		55	2			25		<1									Beryl. 1	crystal.	1/4
1004	40				layer.	1/8	50		35	1			15		<1									1		0.5
2004	do.	do.	Grosscutting.	-	Hanging wall layer Footwall layer.	3/4	30 68		10	2			50		5	1/4	<1			900						

Aumber and	Wall ro	olc												Pe	gmatit	•		100000	4						
ame of peg-		Alteration	Relation to	Shape		Texture										Я	Eneral	ET							
(Pl. II)	formation		wall rock		structure(inches)	Plagic	class	Pert	hite	Grapi	hic ite	Quarts	Mus	covite	Ga	rnet	Tourme	line	Lithium	minera	10	Othe	r mineral	•
							Per-						er- Siz							Mineral		Size inches)	Mineral	Per- cent	Sise (inche
1004 (Continued	Granite and horn- blende gneiss.	None.	Crossoutting	Lenticular	. Core.	5			30				70			<1							Beryl. 5	prystale	1/16-3
1005	do.	do.		Lenticular branching.	- One unit.	1/5-1/	65		15				50	<1		Trace							do. 1	orystal.	1/4
1006	Hornblende gneis		Not exposed.	do.	do.	1/8	58		20	5			20	5	1/8								40. 9	ystals.	1/32-1
1007	do.		do.	Lenticular	. do.	1/8-1/4	54		20				25	1		Trace					350				
1006	Hornblende gneise and fine-grained granite.		do.	Lenticular branching.	- do.	1/8-1/4	55		15				30	<1		Trace									
1009	do.	Mone.		Lenticular	. do.	1/4-1/2	47		33				20	Trac		Trace							Beryl. 2	crystal	.1/4-1
1010	do.		Not exposed.	Lenticular- branching.	do.	3/16	55		30	5	THE .		15	Trace	1/16								40. 6	crystals	.1/16-
1011	Hornblende gneis ine-grained grant and coarse-grained granite.	to.		40.	do.	1/4	45		35				20	<1		Trace							do. 1	crystal.	1/4
1012	Hornblende gneise and fine-grained granite.		Not exposed,	Lenticular	. do.	1/4	50		30				20	<1		Trace							do. 2	rystals.	1/4
1013	Hornblende gneise and coarse-graine granite.			Lenticular branching.	- do.	1/4	50		29				20	1		Trace									
1014	Hornblende gneice fine-grained granite, and coarse-grained granite.	, do.	Not exposed.	Lenticular	. do.	1/5-1/4	55		25				20	<1		Trace							Beryl. 2	crystals	. 3/4
1015	do.	do.		do.	do.	1/8-1/4	60		20				20	<1		Trace							do. 5	crystal s	.1/4-1
1016	Hornblende gneis		Not exposed.	do.	do.	3/16	53		25	3			20	2	1/8	Trace	1/12						do. 1	crystal.	1/8
1017	do.		do.	Lenticular branching.	- do.	1/8	57.5		20	2			20	5	1/8	0.5	1/64								
1018	do.		do.	Lenticular	. do.	1/8	45		30	3			20	5	1/4	0.1	1/16							120	
	Hornblende gneisend fine-grained granite.		do.	Lenticular branching		1/8	142		30	2	50	3	15	8	1/8	Trace	1/64								
1020	40.	None.		Lenticular	. do.	1/4	30	To the last	40	3	30	3	20	10			43.5					SIT-			
1021	do.		Not exposed.	do.	do.	1/8	50	4	25	3	10		20	5	1/8	Trace	1/32		1						
1022	Hornblende gneis	•	do.	do.	Wall sone.	1/16	50 10	1/32		3			28 30	2									Biotite.	Trace.	
1023	do.		do.	do.	Wall some.	1/8	60		19	3			20	1											

0	2
100	- A
E	0
-	

Munber and	Wall r	ock							1000	-200					Peg	matite	0									
name of peg-	Type and formation	Alteration	Relation to	Shape		Texture											M	neralo	ey							
(Pl. II)	IOTMATION		Wall rock		structure(inches	Plagi	oclase	Pert	hite	Gray		Qua	rtz	Muso	ovite	Gaz	net	Tour	maline	Lithiu	n miner	als	Other	r mineral	
							Percent	Size	Per- cent(Size inches	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Size inches	Per-	Size inches	Mineral	Per-	Size (inches)	Mineral	Per- cent	Size (inches
1024	Hornblende gneis	•	Not exposed.	Lenticular		1/2	2 20		28 55	2-1/	2		30 25		2											
1025	do.		do.	do.	Wall zone. Core.	1/8	55 30		10 37				30 30		5 3											
	Fine-grained granite and horn blende gneiss.	Mone.		do.	Wall sone. Core.	3/8	45		25 20	5	15		80		10	3/8	Trace.	1/32								
1027	do.	do.		do.	Wall sone. Core.	1/16	60		20 35	3			15 45		5	1/2	Trace.	1/16						Beryl. 2	crystals	1/4-3/
	Hornblende gneis fine-grained gra- and coarse-grain- granite.	nite		Lenticular branching		1/8	63		10 30	1 4			15 50		5 5	1/8	2	1/64					10	Beryl. 1		
1029	Hornblende gneiss and fine-grained granite.	do.		do.	One unit.	1/2	40		30	4			20		10	1/2	Trace.	1/16								
1030	do.	do.		do.	do.	3/4	35		35	5			20		10	1/2	Trace.	1/16								
1031	do.		Not exposed.	Lenticular	do.	1/8	49		30	3		1	20		Trace.	.1/16	1	1/64				23.0		Beryl. 1	crystal.	3/16
1032	do.		do.	do.	do.	1/4	53		25	4			20		2	1/2	Trace.	1/16							crystal.	
1033	Hornblende gneis	. 6 2 5 7	do.	do.	do.	3/8	49		25	4			20		5	1/4	1	1/64		100					crystals	
1034	do,		do.	Lenticular branching	do.	3/16	39.5		30	3			20		10	1/4	0.5	1/64							crystal.	
	Hornblende gneise and fine-grained granite.		do.	Lenticular	do.	1/16	61		20	3			15		2	1/8	2	1/64						Columbite-	1 cryst.	1. 3/8
1036	do.		do.	do.	do.	1/4	43.5		25				20		10	1/2	0.5	1/128						Cleaveland ite. Biotite. Columbite- tantalite.	1 Trace.	
1037	do.		do.	do.	do.	3/8	34.5		35				20		10	1/8	0.5	1/128						Beryl. 2	erystals.	1/16-1
	Hornblende gneiss fine-grained gra and coarse-graine granite.	nite.		Lenticular- branching.	do.	1/2	40		30	3			20				Trace.									
1039	Hornblende gneiss and fine-grained granite.		Not exposed.	Lenticular	. do.	1/8	60		15	3			20		5	1/4	Trace	.1/128						Beryl. Biotite.	l crysts	1. 3/8
	Hornblende gneis fine-grained gran and coarse-graine	ite.		enticular- branching.	Hanging wall layer	1/4	40		40	4			15		5	1/4	Trace.	1/16								-
	granite.				layer.	1/32	78		5	1/2			15		Trace.	1/16	2	1/128						Beryl. 1	crystal.	3/8

Mumber and	Wall r	OCK						Un Page				-	100		Peg	matite								VI SA	Control of the
name of peg-	Type and formation	Alteration	Relation to	Shape		Texture	HOT										M	neralo	EY						
(Pl. II)	Tornacton		WALL FOCK		structure	inches		oclase	Pert	hite	Gray		Qua	arts	Musc	ovite	Gar	net	Tour	maline	Lithium	minerals	Other	nineral	•
				Ta te			Percent	Size	Per- cent(Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size inches	Per-	Size (inches)	Mineral	Per- Size cent(inches	Mineral	Per- cent	Sise (inches
1041	Hornblende gneis and fine-grained granite.	Hone.		Lenticular branching		3/8	42		35	4			20		3	1/8	Trace.	1/64					Biotite. Chlorite. Beryl. 4	Trace. Trace. crystal	1/4 1/8 •.1/16-1
1042	do.	do.		do.	do.	1/8	58.5		20	3			20		1	3/16	0.5	1/64						crystal	
1043	do.		Not exposed.	Lenticular	wall layer Footwall	3700	53-9		35				15		0.5	1/8	frace.	1/64					Biotite. Beryl.	Trace.	1/4
1044	W	2.7			layer.	1/128	85						15				<1						do. 1	crystal	1/8
1044	Hornblende gneis		do.	branching	. Core.	3/8	78 35		35	3	5		15		10	1/2	2 Trace.	1/32					Beryl. 3 Biotite.	trace.	1/8-3/1
1045	do.	11.	do.	Lenticular	One unit.	1/16	58		20	1			15		5	1/8	2	1/12	8				Beryl. 3	erystals	3/8-1/8
1046	Hornblende gneis and fine-grained granite.		do.	do.	do.	1/16	62		15	3			20		3	1/4	<1	1/32					40. 3	erystale	1/16-3/
1047	Hornblende gneis		Not exposed.	Lenticular branching.	- One unit	. 1/4	49		30	4			50		1	3/16	frace.	1/32					Beryl. 2	erystals	1/8-1/4
1048	Hornblende gneise and fine-grained granite.	None.		Lenticular	do.	1/16	57		20	4			20		3	1/8	frace.	1/32					do. 1	crystal.	1/8
1049	Hornblende gneis		Not exposed.	do.	do.	1/16	56		20	3			20		4	1/8	race.	1/32					Columbite-	1 cryst	1. 3/16
1050	do.		do.	do.	do.	1/8	49.5		30	3			20		0.5	1/2	<1	1/32					Biotite.	Trace.	1/8
1051	do.		do.	do.	do.	1/8	45		35	2			15		5	1/8		1/64							1,0
1052	Hornblende gneiss and fine-grained granite.	None.		Lenticular- branching.	do.	1/4	45		30	14			20		5	1/8		1/16					Beryl. 1	erystal.	1/4
1053	Fine-grained granite.	do.		Lenticular.	do.	1/8	49.5		30	3			20		0.5	1/16	Trace.	. 1/64					do. 5	crystals	3/32-3/
1054	Hornblande gnei		Not exposed.	do.	do.	1/4	43		30	3			20		7	3/16	Trace	1/39					Editor Park	MAG (B)	11. 30. 50.50
1055	do.		do.	do.	do.	1/8	55		20	3			20		5	1/4		1/64	2364				do.	Trace.	1/4
1056	do.		do.	do.	do.	1/4	45		30				25		<1	-/-	-1	1/64							
1057	Fornblende gneiss fine-grained gran and quartzite.			Lenticular- branching.	do.	1/4-1/2	TO SUIT		种				25		1,		Trace								
1058	Fine-grained granite.	do.		do.	do.	1/4	49		30				20		1		Trace								
1059	Mornblende gneiss and quartzite.	do.		do.	do.	1/4-1/2	43		37				20		<1		Trace								
1060	Hornblende gneis		Not exposed.	do.	do.	1/8-1/4	60		15				25		Trace.		<1								

Number and	Wall r	OGK													reg	mat1te										
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture											Miner	alog	y							
(P1. II)	1011120101		wall rock		· ·		Plagi	oclase	Pert	hite	Grap	hic ite	Quar	rtz	Muso	ovite	Garnet		Tourn	aline	Lithium	minera	als	Other	mineral	•
							Percent		Per-			Size		Size	Per- bent	Size	Per- Si bent in	ze l	Per-	Size	Mineral		Size	Mineral	Per- cent	Size (inches
1061	Hornblende gnei	··.	Not exposed.	Lenticular	. One unit	1/8-1/1	55		25				20		<1		Trace.									
1062	do.		do.	do.	do.	1/4	45		35				20		<1		Trace.							Magneti te	Trace.	
1063	do.		do.	do.	do.	1/2	30		50				20		<1		Trace.				2 10 20 10			Biotite.	Trace.	
1064	Hornblende gnei fine-grained grand quartite.	nite,		do.	do.	1/4	49		25				25		1		Trace.							Beryl. 1 Columbite tantalite.	80000	
1065	Hornblende gnei	·.	Not exposed.	Lenticular- branching.	do.	1/4-1/2	45		35				20		<1		Trace.							Beryl. 1	crystal.	1-1/4
	Hornblende gneis and fine-grained granite.			do.	do.	1/2	42		37				20		1		Trace.									
	Hornblende gneise fine-grained gram and quarzite.	do.		Lenticular	. do.	1/8-1/4	60		20				50		<1		Trace.									
1068	Hornblende gneis		Not exposed.	Lenticular branching.	Core.	.1/8-1/ 3	51		25 40				20 58		1		<1							Beryl. 17 do. 3 Columbite- tantalite.	crystal	
	Fornblende gneis and fine-grained granite.			Lenticular	. One unit	1/4-1/2	112		37				20		1		<1							Beryl. 4	in the second	
1070	Hornblende gneis	•.	Not exposed.	Irregular.	do.	1/8-1/4	69		10				20		1		<1							Beryl. 5	crystal	.1/16-
1071	do.		do.	Lenticular	.Wall zone.	1/2	40		35 38				25 60		<1 1		_<1								crystale	in equity
1072	do.		do.	do.	One unit.	/2-3/4	. 35		40				25		<1		Trace.							Beryl. 26	crystal	.1/16-1
	Hornblende gneis and fine-grained granite.			do.	do.	1/4-1/2	##		35				20		1		Trace.									
1074	Hornblende gneis	••	Not exposed.	Lenticular ranching.	- do.	3/4	35		45				20		Trace.		Trace.							Biotite.	Trace.	
1075	Hornblende gneis and coarse-grain granite.	s None.		do.	do.	1/4	45		34				20	01	1		<1							Beryl. 3	crystal	. 3/16
1076	Hornblande gneis and fine-grained granite.		Not exposed.	Lenticular	.Wall zone.	1/2	30 10		50 50				20		Trace.		Trace.							Biotite. Martite. Beryl. 2	Trace. Trace.	.1/16-1
1077	Hornblende gneis	•.	do.	Lenticular branching.	- Wall sone	.1/4-1/	2 45		35 7				20 93		<1		Trace.							Biotite.	Trace.	
1078	do.		do.	do.	One unit.	1/8-1/4	50		30				20		<1		Trace.						5			
1079	do.		do.	Lenticular	. do.	1/4	45		30		WAS S		25		<1		Trace.									1391

Mumber and	Wall r	ock				-									Peg	gmatite										
name of peg-	Type and	Alteration	Relation to	Shape		Texture											M	neral	ogy							
(Pl.II)	formation		wall rock		structure	(inches)	Plagio	clase	Perth	1te	Grapi		Qua	urts	Musc	covite	Gar	net	Tours	naline	Lithium	minerals		Other	mineral	
							Per-	Sise inches	Per-	Size	Per-	Size inches	Per-	Size	Per-	Size	Per-	Sise	Per-	Size inches	Mineral	Per- Sis		lineral	Per- cent	Size
1080	Hornblende gneis and fine-grained granite.	Mone.		Lenticular branching.	- One unit.	.1/8-1/4	55		25				20		<1		Trace.									
1081	do.	do.		do.	do.	1/8	54		25				20		1	150	<1		1				Bei	ryl. 1	crystal.	1/8
1082	do.	do.		do.	do.	1/4-1/2	49		30				20		1		<1						Max		Trace.	-,-
1083	do.	do.		Lenticular	. do.	1/4	60		20			-	50		<1		<1			No.			Bet	ry1. 6	crystals	1/8-3/
1084	do.	do.		do.	do.	1/8-1/4	59		20				20		1		<1			100				do. 5	crystals	1/8-1/
1085	Hornblende gneis		Not exposed.	do.	do.	1/2-3/4	43		30				25		2		Trace.								crystals	
1086	do.			Lenticular branching.	- do.	1/8-1/4	149		30				20		1		Trace.							do. 1		
1087	do.		do.	do.	do.	1/8-1/4	60		20				20		<1		Trace.									
1088	Hornblende gneis and fine-grained granite.	Hone.		do.	do.	1/4	1414		35				50		1		<1							otite.	frace. crystals	3/16
1089	Hornblende gneis		Not exposed.	Lenticular	do.	1/8-1/4	49		30				20	De e	1		<1									
1090	do.		do.	do.	do.	1/4	55		25				20		<1	1 57	<1						24	otite.		
1091	do.		do.	do.	do.	1/2	39		40				20		1		frace.								Trace.	- 0
1092	do.		do.	do. 1	Wall some.	1/8-1/4	55		25				20 74		<1 <1		<1							ryl. 1 do. 1		1/4
1093	do.		do.	do.	One unit.	1/4-1/2	40		35				25		<1		frace.									
1094	do.		do.	do.	do.	1/2	35	110000	40				25		<1	1	Trace.								erystals	
1095	do.		do.	do.	do.	1/4	45		30				25		<1	-			DAG						erystals	
1096	do.		do.	do.	Wall sone. Core.	1/2	35		45 25				20 74		<1		race.						•	do. 2	erystals	1/8
1097	do.		do.	do.	One unit.	3/4	15		65		15		20		<1											
1098	do.		do.	do.	do.	1/4	hh		35	8.0	-,		20		1	DW 18	race.			- 199					The same	
1099	do.		do.	do.	do.	1/2-3/4	40		39				50		1	-3300	race.									
1100	do.		do.	Lenticular- branching.	do.	1/4	1614		30				25		1		race.						Ber	ryl. 2 c	rystals.	1/4
1101	do.		do.	Lenticular	. do.	1/4	45		30				25		<1		Trace		Libbs.							
1102	do.		do.	do.	do.	1/4	49	1	30				20				1000							idote.	Trace.	
1103	do.		do.	Lenticular branching.	X	1/4	##		35				20		1		Trace						Ber	ryl. 1	rystal.	3/16
1104	do.		do.	Lenticular	. do.	1/8	46.5		30		6	10- 11	20		3		0.5									

umber and	Wall r	ock					1			21002	477.1				Peg	matite										
ame of peg-	Type and formation	Alteration	Relation to	Shape		Texture											M	neral	ogy							
(P1. II)	Tormation		ANTI LOCK		structure	inches	Plagi	oclase	Pert	hite	Gray	hic ite	Que	rts	Musc	ovite	Gaz	net	Tour	naline	Lithium	minerals		Other	mineral	•
							Percent	Size	Per-	Size				Size					Per-	Size	Mineral	Per- S cent (in	ize iches)	Mineral	Per- cent	Size (inches
1105	Hornblende gneis		Not exposed.		wall layer. Footwall	3/4	30		45	4			15		10	3/4	Trace	1/64								
1106	do.		do.	do.	layer.	1/16	67.5		15	2			15		2	1/4		1/128						Beryl. 1	crystal.	1/1
Maria San	Fine-grained	Mone.		do.	Hanging	1/4	50		35	3			15		1 5	1/8	Trace	1/64								
	granite and horn blende gneiss.				Footwall layer.	1/32	72		10	1/2																
1108	do.	do.		Lenticular		1/8	44.5		30	2			15		1 5	1/16	0.5	1/64								
1109	do.	do.		branching.	do.	1/16	73		10																-	
1110	do.	do.		Lenticular,		1/16	68		15	1			15		5	1/4										
ıııı	do.	do.		Lenticular branching.	do.	1/8	45		30	2			20		5	1/2										
112	Hornblende gneis		Not exposed.	do.	do.	1/8	50		25	1	15	h	15		2	1/8										
1113	Fornblende gneise and fine-grained granite.	Mone.		Lenticular	do.	1/2	39		45	3			15		1		Trace.	1/64								
1114	do.	do.		Lenticular- branching.	do.	1/32	74		10	3/4			15		1	1/16	~1	1/64								
1115	do.	do.		Lenticular	do.	1/16	72		10				15		3	1/8										
1116	do.	do.			Hanging vall layer. Footwall	3/8	20		60	3/4			15		5	1/4										
					layer.	1/64	84						15		1	1/64										
117	Hornblende gneiss		Not exposed.		Hanging wall layer Footwall	1/2	15		140	1-1/2			40		5	1/8										
118	Hornblende gneise and fine-grained granite.	None.		Lenticular branching.	layer.	1/64	78		5				15 15		Trace 2		<1 Trace.	1/64	,							
119	Hornblende gneise		Not exposed.		wall layer. Footwall	3/8	8		60	3/4			25		7	1/4										
120	do.		do.	Oval.	layer.	1/32	83						15		2	1/64										
121	do.		do.	Lenticular	One unit.	3/8	39 59		20	1-1/2			15 20		1	1/4	<1	1/8								
122	do.		do.	branching.	Hanging	1/8	20		60																	
				TO LOCAL DE LA SE	wall layer.	1/16	20		a				19		1											

1	N	2
	D	2
	1	4

Aumber and	Wall r	ock													Peg	mat1te			15							
same of peg-	Type and	Alteration	Relation to	Shape		Texture (inches)											M	neralo	EY							
(Pl. II)	formation		wall rock		structure		Plagic	clase	Perti	hite	Grap	hio ite	Quar	ts	Muso	ovite	Gar	net	Tours	aline	Lithium	minera	J.	Other	mineral	10
							Per-	Size (inches	Per-	Size		Size Poinches	nt i	Size nches	Per-	Size	Per-	Size	Per-	Size inches)	Mineral	Per-	Size inches	Mineral	Per- cent	Siscinches
1123	Hornblende gnei	•	Not exposed.	Lenticular	Hanging wall layer Footwall	1/4	60		10				27		3											
				do.	layer.	1/64	75		5 24	h			30		2											THE STATE OF
1124	do.		do.	do.	Hanging	1/4	17		50				30		3	3/4										
1125	do.		ao.	ao.	wall hayer Footwall layer.	1/64	60		19				20		1) -										
1126	do.		do.	do.	Hanging	1/2	15		50				30		5	1										
					wall layer Footwall layer.	1/32	60		8				30		2		-									
1127	do.		do.	do.	One unit.	3/8	30		30				30	,	10											
1128	do.		do.	do.	do.	1/16	70		14				25		1		- 3 10									
1129	do.		do.	do.	Hanging wall layer	1/8	10		60				27		3											
					Footwall layer.	1/32	60		-/				38		2											
1130	do.		do.	do.	Hanging wall layer Footwall	1/4	30		50				28		2											
					layer.	1/32	70		14				25		1											
1131	do.		do.	do.	Hanging wall layer Footwall	1/8	10		60				27		3											
					layer.	1/32	70		74				25		1							F. Co				
1132	do.		do.		Hanging wall layer Footwall	TO MENT HE	10		50				40		<1											
1127			40	40.	layer.	1/32	80		10				25		5							7.7				
1133	do.		do.	40.	Core.	1/32	30		50				49		1											
1134	do.		do.	Oval.	One unit.	1/16	60		10				30													
1135	do.		do.	Lenticular	. do.	1/8	50		50				30		Frace											
1136	do.		do.	do.	do.	1/8	₩.5		30	5			15		0.5	1/8	<1	1/84								
1137	Hgrnblende gneis and fine-grained granite.	s Hone.	do.	Irregular.	One unit.	1/4	48		30				20		2	1/4										
1138	do.	do.	do.	Lenticular	. do.	1/8	60		20	5			20				Trace	1/64				HA C		Biotite.	Trace.	1/
1139	Hornblende gneis		do.	Irregular.	do.	1/16	75		10	3			15		182											
1140	do.		do.	do.	do.	1/4	58		25	4			15		2	3/8										
1141	do.		do.	Oval.	do.	1/4	75		10	2	A CONTRACTOR		15	O SH	20	BIN		I B Tree				1	ABB			

Quarts

Graphic granite Pogmatite

Muscovite

Mineralogy

Tourmaline

Lithium minerals

Other minerals

Garnet

Wall rock

Alteration

Relation to

wall rock

do.

Lenticularbranching.

Shape

Type and formation

Number and name of pegmatite (Pl. II)

1159

1160

do. .

do.

	NEW CONTRACTOR OF THE PARTY OF		9							1000	- Rren	11 00						1		3 11 11						
							Per- cent	Size (inches	Per-	Size inches	Per-	Size	Per-	Size (inches	Per-	Size	Per-	Sise	Per-	Size inches	Mineral	Per-	Size inches	Mineral	Per- cent	Size (inche
1142	Hornblende gneis		Not exposed.	Lenticular	. One unit	1/2	50		35	5			15													
1143	do.		do.	Oval.	do.	1-1/2	45		35		15	8	20													
1144	do.		do.	Irregular.	do.	1/4	75		10	2			15	1	5	1/4										
1145	do.		do.	Lenticular- branching.	do.	1/4	50		25		15	6	15		10	1/4				- 51						
1146	do.		do.	Lenticular	. do.	1/8	50		20	3	10	5	50		10	1/4										
1147	do.		do.	Lenticular branching.	- do.	1/4	56		25		15	6	15		3	1/8	1	1/128								
1148	do.		do.	do.	Wall mone.	1/8 3/4	58.5		25	2 1/2			15 30		1 55	1/8	0.5	1/64								
1149	do.		do.	Lenticular	. One unit.	1/16	72		10	2			15		3	1/4										
1150	do.		do.	do.	Wall sone. Core.	1/8	10		30 30	3			20 55		10 5	1/4								Beryl.	2 crystal	.1-1/2-
1151	Coarse-grained granite and hornblende gneiss	None.		do.	One unit.	1/8	45		35	3			15		5	1/4	Trace	1/12								3-1/4
1152	do.	do.		do.	Hanging wall layer. Footwall	1/8	40		40	3			15		5	1/4										
					layer.	1/64		246	84				15				1	1/128								
1153	Hornblende gneis		Not exposed	branching.	wall layer		40		40	1			15		5	3/8	Trace.									
					layer.	1/64	84						15				1	1/128								
1154	Fine-grained granite and horn- blende gneiss.	None.		Lenticular	wall layer Footwall	Day Inchis	140			1-1/2			15		1	1/2		1/128								
					layer.	1/64	80		5	1/2			15				Trace.	1/128								
1155	Hornblende gneiss		Not exposed	Lenticular branching.	Hanging wall layer Footwall		25		45	5			20		10	3/4										
					layer.	1/64	79		5	1/4			15				1	1/128								
1156	do.		do.	Lenticular	Hanging wall layer Footwall	the State	48.5		35	2	25	5	15		1	1/4	0.5	1/128								
					layer.	1/64	84						15				1	1/128								
1157	do.		do.	Lenticular branching.	One unit.	1/2	15		50		40	5	20		15	1/4	<1	1/32								
1158	do.		do.	Lenticular	. do.	1/4	55		20	5	B) S	1000	25	1		FILE										

20

15

Trace. 1/16

Trace. 1/8

Perthite

Internal Fexture structure (inches)

1/8

1/8

65

59

15

25

are a 710

Mumber and	Wall r	ock													Peg	mat1te	•		1	315						
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture											N	neral	DEY.							
(P1.II)	10Finaction		Wall rock		structure	Inches	Plagi	oclase	Pert	hite	Gray	phic nite	Qua	rts	Musc	ovite	Gaz	rnet	Tour	maline	Lithiu	m miner	ale	Other	mineral	
							Percent						Per-						Per-	Size	Mineral	Percent		Mineral	Per-	Size (inches
1161	Hornblende gnei	••.	Not exposed.	Lenticula	r. One unit	1/4	65		10	4			25						1780							
1162	do.		do.	do.	do.	3/8	60		10	1-1/2			30													
1163	do.		do.	do.	do.	1/8	70	10.45	15	1			15				Trace	1/6								
1164	do.		do.	Lenticular- branching	Hanging wall some	1/2	30		50	1			20		5	1/4										
					layer.	1/64	75		10	3/4			15		Trace	1/32.	Trace.	1/128								
1165	do.		do.	Lenticular	.Wall zone Core.	1/16	60 15		20 38	2			15 45		5 2	1/16										
1166	Hornblende gneis	s None.	do.	do.	One unit.	1/8	62	7.57	50	2			15		3	1/4	Trace	1/128								
1167	Fine-grained granite and horn blende gneiss.	do.		do.	do.	1/8	50		25	3		500	20		5	1/4									,	
1168	do.	do.		do.	do.	1/16	69		10	1			20		1	1/16				6013						
1169	do.	do.		Lenticular branching.	- Wall son		53		25 10	1 4	15	5	20 80		2 10	1/8										
1170	Fornblande gnets		Not exposed.	do.	One unit.	1/8	50		20	3			20		10	1/4										
1171	Hornblende gneis and fine-grained granite.			Lenticular	. do.	1/2	30		50	8			15		5	3/8.	<1	1/64								
1172	do.	do.		Lenticular branching	- Hanging .wall layer	1/4	35		40	5			15		10	1/4	<1	1/64								
					Pootwall layer. Core.	1/64	75		45	5			15 45		10 10	1/64	Trace							Beryl. 1	crystal	1
1173	Hornblende gnets	. 1. 67	Not exposed	Lenticular	. One unit.	1/8	69.5		15	2			15		0.5	1/4	<1	1/32								
1174	do.		do.	Lenticular branching.	-Wall sone. Core.	1/8	60 7		20 50				20.		<1 <1		Trace							Beryl. 5	rystals.	. 1/4-1
1175	do.		do.	Lenticular	One unit.	1/4-1/2	54		25				20		1		Trace									
1176	do.		do.	do.	do.	1/4-1/2	1414		35	I PLEAS			20		1				100		No.					
1177	do.		do.	Irregular.	Wall zone.	1/2	53		25 30				20		2 3		<1									
1178	do.		do.	do.	Wall sone.	1/2	35		41 15				20 85		47		<1									
1179	do.		do.	do.	One unit.		40		40		- 10		19		1		<1									
1180	do.				filling.	3	2		58		S.RS		40		<1									Beryl. 8	crystal	.1/4-1/
			do.	Lenticular			196		50		G S		30		1		<1	1	Dist.							
1181	Hornblende gneis and fine-grained granite.			Lenticular		1/8-1/4	45		33				20		5		Trace		-			1				

humber and	Wall ro	ck					5000	3 100			100		NG-		Peg	mat1te			213						1000	0.54
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture		14.7									M	neralo	EY							
(P1. II)	10Fmat10H		ANTI LOCK		structure		Plagio	class	Pert	hite	Grap		Que	rts	Musc	ovite	Gar	net	Tours	naline	Li thium	minera	1.	Other	mineral	•
							Per-					Size								Size inches)	Mineral	Per-	Size inches)	Mineral	Per- cent	Si:
1182	Hornblende gneis		Not exposed.	Lenticular branching.	- Wall sone	.1/8-1/	4 40		40 5				20		< 1									Beryl. 2	crystale	. 1-6
1183	do.		do.	Irregular.	One unit.	3/4	25		45				30				<1									
1184	do.		do.	Lenticular branching.	- do.	1/4	43		. 35				20		5		Trace.									
1185	Hornblende gneise and fine-grained granite.	Fone.	do.	do.	do.	1/4	43		35		Trace.		20		2		Trace.									
1186	Hornblende gneis		Not exposed.	Lenticular	. do.	1/4-1/2	15		57				25		3		∠1									
1157	do.		do.	do.	Wall sone. Core.	1/2 6-8	10		68 15		5		20 85		2		<1									
1188	Hornblende gneiss and fine-grained granite.	None.		Irregular.	Wall sone. Intermedia sone. Core.		62 35		30				15 25 100		3 20		l Trace.									
1189	Hornblende gneiss		Not exposed.	Lenticular	Wall zone Core.	.1/4-1/2	44 <1		35				20		1 <1		<1							Beryl. 1 c	rystal.	3/
1190	Coarse-grained granite and horn- blende gneiss.	None.		do.	One unit.	1/2	30,		50				20		<1		Trace.			4						,
1191	Hornblende gneis		Not exposed.	Lenticular branching.	- do.	1/2	42		37				20		1		Trace.									
1192	do.		do.	do.	do.	1/2	35		111				20		1		Trace.							Beryl. 22	crystals	1/8-
1193	Hornblende gneiss fine-grained gram and coarse-graine	ite,		do.	Wall sone.		50		20	5			20		10	1/2	Trace.	1/64								
	granite.				Core.	8	15		25	5	10		35 80											Hematite. Beryl, 3	Trace.	1-1-
1194	Hornblende gneiss		do.	Lenticular,	One unit.	3/16	48		30				20		2		Trace.								crystals	
1195	do.		do.	do.	do.	1/8	42		30	3			20		3	1/8	Trace.	1/32							· -	
1196	do.		do.	do.	Wall zone. Core.	1/8	59-5		20 60	3 8			20		0.5	1/8	Trace.	1/64								
1197	Hornblende gneiss and fine-grained granite.	None.		do.	Vall zone. Core.	1/8	60		25	2 8			15 60		Trace.		Trace	1/64						Biotite.	Trace.	1/2
1198	do.	do.		do.	Wall zone. Core.	1/8	59 10		20 50	2 5			20 35		0.5	1/4 1/2	0.5 Trace.	1/32 1/32							,	
1199	do.	do.		Lenticular branching.			48		30	5			50		2	1/4	<1	1/32								
				- anoning.	sone.	1 8	40		10 10	5			30 90		20	2	Trace.	1/32								
1200	do.	do.		do.	One unit.	1/4	62		15	1		188	20	827 B	3	1/8	Trace.	1/32				78.	A CONTRACTOR			

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Mumber and	Vall r	oak												Po	matite			The Sa	E STO	Market Service					
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture (inches										MI	neral	67							
(Pl. II)	Tormacton		Wall rock.		structure	(Inches		oclase	Pert	thite	Grap	hic ite	Quarts	Muse	povite	Gar	net	Tours	aline	Lithiu	miner	als.	Other	mineral	•
							Percent	Size	Per-	Size Inches	Per-	Sise Pe	nt inc	Per-	Sise	Per-	Sise		Sise inches	Mineral	Per-		Mineral	Per-	Size (inches
1201	Hornblende gnei		Not exposed.	Irregular.	One unit.	3/8	39		40				20	1	1/4	<1	1/64						Beryl. 1 c	rystal.	1/2
1202	do.		do.	Lenticular	.Wall some Core.	1/4	50 5		30 35	1-1/2			80 60	Trace	1/4	Trace.	1/64						do. Beryl. 4 c	Trace.	1/4
1203	do.		do.	do.	One unit.	1/4	140		35	1			5	Trace	1/4	Trace.	1/64								
1204	do.		do.	do.	do.	1/16	56.5		20	3			10	3	1/8	0.5	1/64						Magnetite.	Trace.	1/4
1205	do.		do.	do.	do.	1/16	55		20	3/4			10	5	1/8	<1			NEW S						
1206	do.		do.	Lenticular			50		25	3			0	5	1/8	<1	1/64								32.15
				oranghing.	Intermediat sone. Core.	1/4	46		10	5		2 6	0	20	1	2	1/4								
1207	do.		do.	Lenticular	One unit.	3/16	40		30	5		8	0	10	1/4	<1	1/64								
1208	Hornblende gneis	None.		do.	Wall zone.	1/8	44.5		30.	5		8	0	5	1/8	0.5	1/32								
	granite.				sone. Core.	3/4	25 20		10	3 8	-	3	5	30	1 1/4										
1209	Hornblende gneis		Not exposed.	do.	Wall sone. Core.	1/8	45		25	3		10	0	10	1/4	Trace.	1/32								
1210	do.		do.	Lenticular branching.	One unit.	1/4	53		25	4		8	0	2	1/4	Trace.	1/16								
1211	do.		do.	Lenticular	Wall some.	6-8	15 2		59 30			8	5	1 < 1		<1									
1515	do.		do.	do.	Wall sone. Core.	1/2	30		40	5		10	0	10	1/2	Trace.	1/32								
1213	Hornblende gneis and fine-grained granite.	None.		de.	One unit.	1/4	60		25		15	5 1	5										Biotite.	Trace.	1/4
1214	Fornblende gneis		Not exposed.	do.	do.	3/8	45		40		20	1	5										Chlorite.	Trace.	1/4
1215	do.		do.	do.	do.	3/8	140		35	4			5										Magnetite.		1/16
1216	Hornblende gneise and granite.		do.	do.	do.	3/8	45		30		3	24				0.5	1/64						do. Hematite.	Trace.	-/1
1217	Fornblende gneise		do.	do.	Wall sone. Core.	1/16	75 45		30	3		2	5										Magnetite.		1/16
1518	do.		do.	do.	One unit.	1/4	40		40		10	2	0										Martite.	Trace.	
1219	do.		do.	do.	Wall sone.	1/2	15		60		10		5										do. Biotite.	<1 Trace.	
1220	do.		do.	Lenticular branching.	Gore. One unit	2-3	15		65				8										Magnetite.		
1557	do.		do.	Lenticular	do.	/4-1/2	35		140			2	5			<1							Martite.	Trace.	

Mumber and	Wall ro	District Selection													Pegma	tite		-							SILL VIS	
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture		Marie 1									Mi.	neralo	СУ							
(Pl. II)							Plagi	oclase	Pert	hite	Grapi gran	hic ite	Quart	s 1	Muscovi	ite	Gara	net	Tourm	aline	Li thium	minera	als.	Other	mineral	
							Percent	Size		Size	Per-)cent(Size :	Per- S	ize P	er- Si ent(in	ize	Per-	Size	Per-	Size inches	Mineral	Per-	Size (inches	Mineral	Per- cent	Size (inches
1222	Quarts monzonite	. None.		Lenticular branching.	- One unit	. 1/8	60		20				50	Tr	ace.		Prace.							Magnetite.	Trace.	· v
1223	do.			do.	do.	1/4	40		40				20				<1							do.	<1	
1224	do.			do.	do.	1/4	55		25				20	Tr	ace.	2	Trace.							do.	frace.	
1225	do.			Lenticular	. do.	3/4	20		55		10		25			2	frace.					4.55		Martite. Biotite.	Trace.	
1226	do.	None.		Lenticular branching.		.3/4-1	10		70		15		20			7	Trace.							Martite. Biotite.	Trace.	
1227	do.			Lenticular	Core. One unit.	3/4	25		50	1210	20		25				TAN							Martite.	Trace.	
1228	do.			Lenticular branching.	- do.	1/2-3/4	25		55				20			7	race.							do. Biotite.	Trace.	
	Quartz monzonite and hornblende gneiss.	None.		do.	do.	1/4-1/2	20		55				25				<1							Magnetite.	<1	
1930	Quartz monzonite.			do.	do.	3/4	50		25				25											do. Biotite.	<1 Trace.	
1231	do.	None.		Irregular.	do.	1/4-1/2	35		45	L.			20		<1	T	race.							do. Magnetite.	Trace.	
1232	Hornblende gneiss		Not exposed.	Lenticular- branching.	do.	1/8-1/4	50		30				20			7	race.							do. Biotite.	Trace.	
1233	Quartz monzonite.			Lenticular	do.	1/4-1/2	30		55				15		<1		<1							Martite.	<1	
1234	do.	None.		Irregular.	Wall zone.	1/2	15		60				25				<1							do. Columbite-	<1	
					Core.							1	.00											tantalite.	10 crys	tals.
1235	do.	do.		do.	One unit.	1/4	50		55				25											Martite.	<1	
	Hornblende gneiss		Not exposed.	Lenticular,	do.	1/8	60		15				25				<1							Magnetite. Biotite.	Trace.	
1237	do.		do.	do.	do.	1/32	88		5	1/2			15											do.	2	1/16
1238	do.		do.	Irregular.	Wall zone. Core.	1/32	80 10		5 30	1			15 60	Ti	race. 1	1/16	Trace.	3/8	Crace.	1-1/2				do. do. Beryl. 3 Samarskite	Trace. Trace. trystals 2 crysta	1/64 1/16 3/4-1-
1239	do.		do.	do.	One unit.	1/64	78		5	5			15					2						Biotite.	2	1/64
1240	do.		do.	do.	do.	1/32	75		10				15											do. Columbite- tantalite.	<1	1/32
1241	do.		do.	Lenticular.	do.	1/2	30		55		35	3	15						Trace.	1/8				Biotite.	<1	1/38
1242	do.		do.	irregular.	Wall zone Core.	1/32	50 60		30				20											do.	Trace.	1/16

Mumber and	Wall r					1807.2	_		1			51.14		3179	Peg	matit	0					0		September 1		
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture											M:	neral	Per							
(Pl. II)			Walt Tout			,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,		oclase	Pert	hite	Graj	phic nite	Qua	rtz	Musc	ovite	Gaz	rnet	Tours	naline	Lithiw	miner	ds.	Other	mineral	•
							Percent	Size		Size				Size inches					Per-		Mineral	Per-	Size	Mineral	Per- cent	Size (inches)
1243	Fornblende gneis	•	Not exposed.	Irregular.	Wall sone.	1/16	75	X.=3	5				20						Trace.	1/8				Biotite.	Trace.	
15/1	do.		do.	Lenticular branching.	-South end. Worth end. Core.	1/2 1/64 2-1/2	60 84 60		20	3			20 15 40				Trace	1/64						Magnetite.	Trace.	1/4 1/64
1245	do.		do.	Lenticular		1/64	84						15											Biotite.	1	1/32
1246	do.		do.	do.	do.	1/8	70		15		5	3	15				<1	1/64								
1247	do.		do.	Irregular.	do.	1/16	80		5	1			15											Biotite. Martite.	< 1 frace.	1/8 1/16
1248	do.		do.	do.	Wall zone. Core.	1/8	65		20		10	8	15 100		Prace.	1/64	<1	1/16						Biotite.	frace.	1/2
1249	do.		do.	Lenticular	One unit.	1/4	70		15	6			15		Trace.	1/4	Trace.	1/64								
1250	do.		do.	do.	do.	1/16	75		15		5	5	10		race.	1/16	<1	1/64								
1251	do.		do.	do.	do.	1/16	80		5	1/2			15				<1	1/64						Biotite.	<1	1/32
1252	do.		do.	do.	do.	1/32	75		10	3			15				<1	1/64								
1253	do.		do.	do.	Wall zone. Intermediat		75		10		5	3	15				Trace.	1/16								
					core.	1	30		50 25	3			20 75		Tace.	1/8								Biotite.	Trace.	1/2
	Coarse-grained granits and horn- blende gneiss.	None.	*	do.	One unit.	1/16	75		10	2			15											do. Magnetite.	<1 Trace.	1/32 1/32
1255	Hornblende gneise		Not exposed.	do.	do.	1/8	55		30	3			15		race.	1/16	Trace.	1/64								
1256	do.		do.	do.	do.	1/4	55		30		20	5	15				<1	1/64						Martite.	<1	1/64
198	Coarse-grained granite and horn- blende gneiss.	None.		Oval.	do.	1/8	65		20		10	5	15		race.	1/16								Biotite.	Trace.	1/16
1258	do.	do.		Irregular.	do.	1/16	70		15	1	12.0		15		race.	1/16	<1	1/64								
1259	do.	do.		do.	do.	1/8	55		30		20	5	15				Trace.	1								
1260	do.	do.		Lenticular- branching.	do.	1/16	70		15	3			15					1/32								
1261	do.	do.		Irregular.	do.	1/8	65		20		10	5	. 15		race.	1/8	Trace.	1/32				Pigg				
1262	Hornblende gneiss		Not exposed.	Lenticular,	Wall zone Core.	1/8	70 75		15 5	1 2			15 20		race.	1/32 1/16	<1 Trace.	1/64								
1263	do.		do.	do.	One unit.	1/8	70		10	1-1/2			20				<1	1/64								
1264	do.		do.	Lenticular- branching.	do.	1/8	50		30	3			50				<1	1/16								
1265	do.		do.	Lenticular.	do.	1/8	50	9	30	5	5		20		700		<1	1/64						Biotite.	Trace.	3/8

humber and name of peg-	Wall ro	Alteration	Relation to	Shape	Internal	Texture			5 10					Que la		matite	The Same	neralo	øv			University of				
matite (Pl. II)	formation		wall rock		structure	inches)	Plagio	clase	Pert	hite	Grap		Que	rts	Musc	ovite		net	Tours	aline	Lithium	minera	10	Other	mineral	
								94	D		gran											Τ_				
							cent	Size inches)cent	inches)cent(inches	cent	inches)cent	inches)cent	inches)cent	inches) Mineral	Per-	Sise inches)	Mineral	Per- cent	Sise (inche
1266	Fine-grained granite and hore blende gneiss.	None.		Lenticular	. One unit	1/8	62		20	5			15		3	1/4	Trace	1/12								
1267	Coarse-grained granite and horn blende gneiss.	do.			Wall sone. Intermediates sone. Core.	1-1/2	33 30		50 20 15	8	30	3	15 30 85		2 20	1/4	Trace	1/16								
1266	Hornblende gneis	1.	Not exposed.	Irregular.	One unit.	1/4	45		40		30	5	15											Biotite.	<1	1/4
1269	Hornblende gneis	None.	Crosscutting	do.	do.	1/8	60		15	3			25				<1	1/32						do.	<1	1/8
1270	Hornblende gneis		Not exposed.	Lenticular	. do.	1/4	35		50				15		<1		<1									
1271	do.		do.	Irregular.	do.	3/4	49		30				20		<1		Trace.									
1272	do.		do.	Lenticular	. do.	1/4	63		15				20		2		<1									
1273	do.		do.	Lenticular branching.	- do.	3/4	49		30				20		1		Trace.									
1274	do.		do.	do.	do.	1/4-1/2	58		20				20		2		<1		Trace.	8000				Martite.	Trace.	
1275	do.		do.	do.	do.	1/8-1/4	65		15				20		<1		<1	100					0.50			
1276	do.		do.	Lenticular	. do.	1/8-1/4	65		15				20		<1		<1									
1277	do.	•	do.	do.	do.	1/4	54		25				20		<1		<1		Trace.					Martite. Biotite.	<1 Trace.	
1278	do.		do.	Lenticular branching.	- do.	1/4	45		35				20		Trace.		Trace.		Trace.					Martite.	<1	
	Hornblende gneiss and fine-grained granite.	None.	1	Lenticular	do.	1/8-1/4	90		5				5											Biotite.	<1	
1260	Hornblende gneiss and quartz mon- sonite.	do.		do.	do.	1/4-1/2	140		35				25		<1		Trace.									
1281	do.	do.		Irregular.	do.	1/2-3/4	35		40				25				Trace.							Martite.	<1	
1282	do.	do.		Lenticular	do.	1/2	15		65		5		20				Trace.							Biotite. Nagnetite.	<1 <1	
1283	do.	do.		Irregular.	do.	1/4	35		45		5		20		Trace		Trace.			1000				do.	Trace.	
1284	Hornblende gneiss		Not exposed.	Lenticular	do.	12-24	20		60		70		20											Biotite. Magnetite.	Trace.	
1285	Hornblende gneise and quartz mon- zonite.	None.		Lenticular branching.	Wall zone	8-12	30		50		60		20				Trace.						W. Land	Samarskite (1) Biotite.	. Trace.	
					Core.	6	1		15				84		<1		Trace.							Magnetite. Biotite. Magnetite.	<1 Trace.	

870 9 FEB

Aumber and	Wall r	ock							Line	7.00		11.9%			Peg	matite										
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture							-				M	neral	PEY							
(P1. II)			Wall Tour		301 (100 110		Plagio	clase	Pert	hite	Grap		Que	artz	Musc	covite	Gar	net	Tours	aline	Lithium	minera	1.	Other	mineral	
							Per-	Size	Per- cent(Size inches	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Size	Per-	Size inches	Mineral	Per- cent(Size Inches)	Mineral	Per- cent	Size (inche
1286	Hornblende gneis and quartz mon- zonite.	None.		Lenticular branching.	- Northeast branch. Southwest branch, wall sone. Southwest branch core.	1/4-1/2 1/4-1/2	15 45		65 35 30		80		20 20 69		<1		<1							Magnetite. Biotite. Magnetite. Biotite.	<1	
1287	Quarts monsonite	do.		do.	One unit.	1/2-3/4	45		35				20		Trace.		<1							Biotite.	<1	
1288	do.	do.		Lenticular	. do.	1/4	20		55				25											do. Martite.	<1 <1	
1289	do.	do.		Lenticular branching.	- do.	1/2	15		65				20											Biotite. Martite.	<1 Trace.	
1290	do.	do.		Lenticular	. do.	1/4	30		45				25				Trace.							Biotite. Martite.	Trace.	
1291	do.	do.		do.	do.	1/8-1/4	38		42				20				Trace.							Biotite. Magnetite.	Trace.	
1292	Hornblende gneis		Not exposed.	do.	do.		49.5		30		85		20	Tille Sel	Frace.		0.5							Biotite.	Trace.	
1293	do.		do.	do.	do.	3/8	60		25		20		15		Trace.		Frace.							do.	Trace.	
1295	do.		do.	Lenticular branching.		3/16	59 53		25		20		50		2	Made 2	Frace.							Magnetite.	Trace.	
1296	do.		do.	Lenticular	do.	1/4	52.5		25		-4-		20		2		0.5									
1297	do.		do.	do.	do.	1/8	48.5		30				20		1		0.5							Martite. Biotite.	Trace.	
1298	do.		do.	Lenticular branching.	do.	3/8	54.5		30		Trass		15		0.5		Prace.									
1299	quarts monsonite.			Irregular.	do.	1	40		45		35	12	15				٠							Magnetite. Biotite.	<1 Trace.	1/16
1300	do.			Lenticular	do.	1/2	55		25		15	5	20				<1	1/32						do.	Trace.	1/4
1301	do.	Mone.		Owal.	do.	5	65		15		5	8	20											do. Magnetite.	Trace.	1/4 1/32
1302	do.	do.		Irregular.	do.	1	55		25		15	8	20									- 2.2		do. Biotita.	Trace.	1/2 1/16
1303	do.	do.		Lenticular,	do.	1/2	55		25	3			20											Magnetite. Biotite.	Trace.	1/2
1304	do.			do.	do.	1/4	60		20	3			20											Magnetite. Biotite.	Trace.	1/4 1/16
1305	do.			do.	do.	1	50		30		20	12	20										Carl I	do.	Traco.	1/32
1306	do.			do.	do.	1/4	65		20		10	8	15									2.00		Magnetite.	Trace.	1/16

hunber and	Vall r	ock												2000	Pog	mat1te										
ame of peg-	Type and	Alteration	Relation to	Shape	Internal	Texture											H	ineral	NS.							
(Pl.II)	formation		wall rock		structure	Inches	Plagi	oclase	Pert	hite	Gray	hic	Que	rts	Musc	ovite	Gas	rnet	Tours	maline	Lithium	miners	g.	Other	mineral	•
							Per-	Size (inche	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Sise	Per-	Size	Mineral		Size inches)	Mineral	Per-	Sis (inche
1307	Quarts monsonite	. None.		Lenticular	. One unit	1/16	75		10	4			15											Magnetite.	<1	1/:
1308	do.			Lenticular branching.	- do.	1/2	40		35	3	5		15											do. Biotite.	<1 Trace.	1/
1309	do.			Lenticular	. do.	1	35		50	3			15											Magneti te.	<1	1/
1310	do.	None.		Irregular.	do.	1/2	20		62				15											do. Biotite.	3<1	1/
1311	do.	do.		Oval.	do.	3/4	45		35	4			20											Magnetite.	<1 Trace.	1/
1312	do.	do.		Lenticular branching.	- do.	2	50		35		25	8	15											Magnetite. Biotite.	Trace.	1/1
1313	do.	do.		Owal.	do.	3	32		50		40	12	15											do. Magnetite.	3 Trace.	1/8
1314	do.			Irregular.	Wall some.	3	40		45		35	12	15 100											do. Biotite.	Trace.	1/1
1315	do.			Lenticular	. Vall sone	. 3	45		40		30		14.5		4									Biotite.	0.5	1
1316	do.	None.		Irregular.	One unit.	4	60		35		25	10	15											Biotite.	<1	2
1317	do.	do.		do.	do.	3	60		25		10	8	15											Martite. Biotite.	<1 <1	1/2
1318	Quarts monsonite			Oval.	do.	4	55		30		20	8	15											Biotite. Chlorite.	Trace.	1/
1319	do.	None.		Irregular.	Zastern part.	5	20		40	6			30		10	1	<1									
					Western part.	6	60		25		15	8	15											Biotite.	Trace.	1/2
1320	do.		No. of the last	Lenticular branching.	One unit.	4	50		30		50	10	19											Magnetite. Biotite.	0.5	1/2
1321	do.			Lenticular	do.	2	60		25		10	8	15											Magnetite. Biotite.	Trace.	1/2
1322	Hornblende gneise and quarts mon- sonite.	None.		Irregular.	Wall zone. Gore. do.	2	55 25		30 40		20	8	15 35 100		10	4	Trace.	1/4	Trace.	2				Biotite. Magnetite. Beryl. 5		1/2-3
1323	Hornblende gneiss		Not exposed.	Lenticular		3	43		40		30		15											Biotite. Magnetite. Chlorite.		
1324	do.		do.	do.	Core.	4	60		35		25		100											Beryl. 3	Trace.	2-1/2
1325	do.		do.	do.	do.	3	65		20		10		15											Magnetite.	Trace.	
1326	do.		do.	do.	do.	2	65		20		2		15											do. Biotite.	Trace.	1/2

Aumber and	Wall r	ock	100												Peg	matite						11:10				
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture			1								M	neral	ogy							
(P1. II)						1011007	Plagi	oclase	Perti	hite	Gran		Qua	rts	Musc	ovite	Gar	net	Tourn	aline	Lithium	minera	18	Other	mineral	
							Per-		Per-	Size nches)	Per-	Size inches	Per-	Size (inches	Per-	Size	Per-	Size (inche	Per-	Size inches)	Mineral	Per- cent (Size inches)	Mineral	Per-	Size (inches
1327	Hornblende gneis		Not exposed.	Lenticular	. One unit	5	47		30		20		20		Trace.	1/16	3	3/8						Martite.	<1	1
1326	do.		do.	Lenticular branching.	- do.	2	59.5		18	4	10	6	20				2					9.79		Magnetite.	0.5	
1329	do.		do.	Lenticular	. do.	6	43		35		25	10	20				Trace.	1/4					3100	do.	2	1/2
1330	do.		do.	Irregular.	do.	1	64.5		20		10	12	15				0.5	1/8						Biotite.	Trace.	1/4
1331	do.		do.	do.	do.	1	59.5		20	5	2		15				5	1/4							0.5	1/2
1332	do.		do.	do.	do.	3	60		25				15											do.	Trace.	
1333	do.		do.		Hanging wall layer. Footwall layer.	5	140		45				15											Biotite.		./-
1334	do.		do.	Lenticular		4	60	0.4	25		15		15				Trace.								1 Trace.	1/8
1335	do.		do.	Oval.	do.	2	144		40		30		15				Trace.	1/8						do.	1	1/8
1336	do.	149.23	do.	Lenticular	. do.	2	53		30		15		15				Trace.	1/16						do.	2	1/4
1337	do.		do.	Lenticular branching.	do.	1	60		25		15		15				. *.	(do.	Trace.	-1/-
1338	do.		do.	Lenticular	. do.	2	50		20				30				<1							do.	Trace.	
1339	do.		do.	Irregular.	do.	1/4	45		35		25		20				Trace.									
1340	do.		do.	Lenticular branching.	do.	3	45		40		30		15											Magnetite.	Trace.	
1341	do.		do.	Lenticular	. Hanging wall layer. Footwall layer.	3/4	35		20		10		20				Trace.							Biotite.	Trace.	
1342	do.		do.	Lenticular			50		30		10	8	20					1/32							Trace.	1/1
				branching.					1									-1,5-								-/-
1343	do.		do.	Lenticular	. do.	2	50		30		50		20											Biotite. Magnetite.	Trace.	
1344	do.		do.	Irregular.	do.	1	65		20			la maria	15											do.	Trace.	
1345	do.		do.	Lenticular	. do.	14	30	P. S	50		5		20				Trace.							Bictite.	Trace.	
1346	do.		do.	do.	do.	5	50		30		10		50											Magnetite.	Trace.	
1347	do.		do.	Irregular.	do.	3	55		30		20	8	15											Biotite. Magnetite.	Trace.	1/2
1348	do.	The State	do.	Owal.	do.	4	60		25		15	12	15											do.	<1	1
1349	do.		do.	Lenticular	do.	1	.20		45		20	8	35		Trace.		Trace.							do.	Trace.	
1350	do.		do.	Irregular.	do.	5	75		10				15		A STATE OF		Trace.							do.	Trace.	
1351	do.		do.	Lenticular	do.	4	45	135	40		20	8	15				Trace.	1/8								3

Aunber and	Wall ro	ock							A STATE OF						Pe	matite									
ame of peg-	Type and	Alteration	Relation to	Shape		Texture											M	neral	ey .						
(Pl. II)	formation		wall rock		structure	(inches)	Plagic	oclase	Pert	hite	Grap		Qua	rts	Muse	covite	Gar	net	Tours	aline	Lithium	minerals	Other	mineral	•
							Per-	Size	Per-	Size	Per-	Sise	Per-	Size	Per-	Size	Per-	Sise	Per-	Size nches)	Mineral	Per- Size cent(inches)	Mineral	Per-	Sise (inche
1352	Hornblende gnei		Not exposed.	Irregular.	One unit.	4	64		20		5	6	15										Magnetite	1	1/2
1353	do.	None.	Crosscutting	Lenticula	. do.	2	60		20		2		20	THE REAL PROPERTY.									Biotite.	Trace.	1-1/4
1354	do.		Not exposed.	do.	do.	4	70		15		5	8	15				Trace	. 3/4					Magnetite.	The second	1/4
1355	do.		do.	Irregular	Wall some Core.	4	50		35	1	25	10	15 100 100				<1			,			Biotite. Magnetite.	<1 <1	
1356	do.		do.	de.	One unit.	ħ	43		40		30	12	15										Biotite.	2	1
1357	Hornblende gneis and quartz mon- sonite.		Crosscutting	. do.	do.	2	60		25		15		15										Magnetite.	<1 Trace.	1/10
1358	Hornblende gneis		do.	do.	do.	1-1/2	75		10	5	1	8	15				<1	1/32					Magnetite,	Trace.	1/8
1359	do.		do.	Lenticular	. do.	1-1/2	55		25		1	4	20										do.	<1	1/4
1360	do.	-	Not exposed.	Oval.	do.	1-1/2	75		10	4	2		15										Magnetite.	Trace.	1/3/
1361	do.	None.	Crosscutting	Lenticular	. do.	1-1/2	55		20		1		25				Trace						Magnetite.	Trace.	1/
1362	Quarts monzonite	. do.		do.	do.	3	45		35		10	6	20				Trace	1/64					Magnetite.	Trace.	1/
1363	Hornblende gneis		Not exposed.	do.	Hanging wall layer. Footwall	2	65		20		10	12	15										Biotite.	Trace.	1-1/
					layer.	3	25		60				15										Magnetite.	Trace.	
1364	do.		do.	One unit.	One unit.	1	74.5		10				15										Biotite. do. Nagnetite.	Trace. 0.5 Trace.	
1365	do.		do.	Oval.	do.	5	55		35		25	8	15										do.	Trace.	1/3
1366	do.		do.	do.	do.	1-1/2	55		30		20	8	15										Biotite.	Trace.	1/3
1367	do.		Crosscutting.	Irregular.	do.	2	65		20		5	12	15										Magnetite.	Trace.	1/3
1368	Quartz monzonite	. do.		Lenticular	. do.	2	40		45		25	10	15										Biotite. Magnetite. Biotite.	Trace.	1/10 3/8
1369	do.	do.		do.	Wall zone.	3	60		25		10	12	14.5									CA SEG	do. Magnetite.	0.5 Trace.	1/8
1370	do.	do.		Lenticular branching.		1	70		15		5		15										do. Biotite.	Trace.	1/3
1371	do.			Lenticular	. do.	2	55		30		20	12	15				Trace	1/4					Magnetite. Biotite.	Trace.	1/1
1372	do.	None.		do.	do.	2	65		25		15	8	15				Trace.	1/4					Magnetite. Biotite.	Trace.	1/1
1373	do.			do.	do.	1	65		20		5	5	15										Magnetite. Biotite. Chlorite.	Trace. Trace.	1/1
1374	do.			do.	do.	1	55		23		2	1000	50				2	1/16					Magnetite.		1/2

235

No.		_	
3	1	2	
100		100	
100	q	9	
13	绉	and .	
470		200	
127	T	7	

humber and	Wall r	ock													Pe	gmat1 te	,								
number and name of peg-	Type and formation	Alteration	Relation to	Shape		Texture											M	neral	ogy				With the second	SELEN	
(Pl. II)	IOTMATION		Wall rock		structure	(inches)	Plagic	clase	Pert	hite	Gran	hic	Qua	artz	Muse	covite	Gar	net	Tour	maline	Li thium	minerals	Other	mineral	•
								Size (inche		Size							Per-			Size	Mineral	Per- Size cent inches	Mineral	Per- cent	Siz
1375	Quartz monzonite	. None.		Lenticular	. One unit	1-1/2	65		20		5		15				Trace	1/8					Magneti te	Trace.	1/
1376	do.	do.		Lenticular branching.	- Wall son	. 2-1/2	30		50		30	12	20				Trace	2					Biotite.	<1	1/
1377	do.	do.		Irregular.	One unit.	3/4	75		10		5		15										Magnetite.		1/
1378	do.	do.		Oval.	do.	1-1/2	65		20		10	8	15				Trace	1/64					Biotite. do. Magnetite.	Trace. Trace.	1/1/1/
1379	do.	do.		Lenticular	. do.	1-1/2	65		20		5		15										do.	Trace.	1/
1380	do.			do.	do.	3/8	69		15				15				1	1/16							
1381	do.			do.	do.	1	70		15		5		15				Trace.	1/16				,	Martite.	<1	1/
1382	do.			Lenticular branching.	do.	1-1/2	60		25	7	5		15				Trace.	1/32					Biotite. Magnetite.	Trace.	1/1/
1383	do.	None.		Irregular.	do.	2	60		25	3	50	8	15										Biotite. Martite.	Trace.	1/
1384	do.	do.		Lenticular	do.	1	70		15		2		15										Magnetite.	Trace.	1/
1385	do.			do.	do.	5	70	5	15		5		15										do.	Trace.	1/
1386	do.	None.		do.	do.	2-1/2	20		55		30		25										Martite.	Trace.	1/
1387	do.	do.		do.	do.	2-1/2	20		55		10		25										do.	Trace.	3/
1388	do.			Irregular.	do.	1	65		20		10	12	15										do. Biotite.	Trace.	1/
1389	do.	do.		do.	do.	3/4	75		10		3	5	15						7				do. Magnetite.	Trace.	1/
1390	do.			do.	Wall sone.	2-1/2	45		35		20	12	20										Biotite.	Trace.	1/
					Core.								100										Magaerite.	mace.	1/
1391	do.			Lenticular	One unit.	1/2	70		15		1		15		Tetra I								do.	frace.	1/
1392	do.			Lenticular branching.	do.	1/2	55		30		5		15										do.	<1	1/
1393	do.			Lenticular	do.	1-1/2	30		55	5	30		15										do. Biotite.	<1 Trace.	3/
1394	do.			do.	do.	3	30		50		40	12	20										do.	<1 Trace.	1,
1395	do.			do.	do.	3/4	60		25		5		15										do.	<1	1/
1396	do.	None.		Lenticular branching.	Wall zone	3	60		25		12		15										do. Biotite.	<1 Trace.	1/
1397	do.	do.			Core.								100		T Ver										
-331		40.		Irregular.	Wall sone	2	50		30		20	15	20		1								Magnetite. Biotite.	<1 Trace.	1/

Number and	Wall ro	ock													Pe	gmatite	•								
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture										19.7	M:	ineral	der.						
(Pl. II)	10Fmation		WAIT FOCK		structure	(inches)	Plagio	clase	Pert	hite	Gran	phic nite	Que	artz	Muse	covite	Gaz	rnet	Tour	naline	Lithium	minerals	Other	mineral	
							Per-					Size (inche		Size		Size				Size (inche	Mineral	Per- Size	Mineral	Per-	Size (inches
1398	Quarts monsonite	. None.		Oval.	One unit.	1/32	10		72				15										Biotite.	3	1/32
1399	do.	do.		Irregular.	do.	1/32	20		65				15										Magnetite.	Z1	1/16
1400	do.	do.		do.	do.	3/4	70		15		10	12	15										Magnetite	Trace.	1/32
1401	do.	do.		do.	do.	1/8	75		10				15										do.	Trace.	1/16
1402 Trio No. 1	do.	do.		do.	Wall zone.	1/4	50		25	1	10	12	15										Biotite. Magnetite.	Trace.	1/16
					Intermedia	6	34	2	30	18			35				0.5	1/2	Trace	ź			Biotite.	0.5	1/2
					do. Core.	1	25	3					60		15	1	<1	1/2					Beryl. 6	crystals	1/2-8
1403	do.	do.		Lenticular branching.	- One uni	. 1/4	55		25				20										Magnetite.	Trace.	1/16
1404	do.			Irregular.	Wall zone.	1/32	60 45		20				20 15										do.	<1 Trace.	1/16
1405	do.			do.	One unit.	1/2	45		15	2			40										do. Biotite.	Trace.	1/16
1406	do.			Oval.	Wall zone.	1/64	79 55		5 30				15 15		Trace	1/64							do.	l Trace.	1/64
1407	do.	None.		Lenticular branching.	-Wall zone.	1/32	79.5		5 30	2			15		Trace.								do.	0.5	.,,
1408	do.	do.		Irregular.	One unit.	3/8	65		15				20										Magnetite.		1/16
1409	do.			do.	do.	1/2	55		30	1			15										Biotite. Magnetite.	Trace.	1/16
1410	do.	None.		Lenticular	do.	1/8	70		15	2			15										Biotite.	<1	1/8
1411	do.	do.		do.	do.	1/2	65		75	2			20										Biotite. Magnetite.	<1 Trace.	1/8
1412	do.	do.		Irregular.	do.	2	30		50		40	8	20										Biotite.	Trace.	1/4
1413	do.	do.		Lenticular	do.	3/4	35		50		40	5	15						A TABLE				Magnetite.	Trace.	1/4
1414 Snowshoe	do.	do.		Irregular.	Wall zone.	. 3	30		50		40		20										do. Biotite.	<1 Trace.	1/4 1/2
1415	do.	do.		do.	One unit.	. 3	30		50		40	8	20										Magnetite. Biotite.	Trace.	1/8 1/2
1416	do.	do.		Lenticular	do.	1	55		15		5		30										Magnetite.	Trace.	1/16
1417	do.	do.		Irregular.	do.	1/2	30		20	3			50										do.	Trace.	1/16
1418	do.	do.		do.	do.	1/4	10		65				25										Martite. Biotite.	Trace.	
1419	do.	do.		Lenticular Core.	Wall sone	3/4	35		45		5	18	20							1			Martite.	Trace.	

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fumber and	Wall re	oek								16.11					Por	gmatite	•	19-2	Au de							1136
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture		1				100					M	neral	ey.							
(P1. II)	Tornaction		WALL FOOK		structure	Inches,	Plagicol	lase	Perti	hite	Gray		Qua	arts	Musc	covite	Gas	rnet	Tours	naline	Li thium	miners	T.	Other	mineral	•
							Per :	Size		Size inches	Per-)cent (Size inches	Per-	Size	Per-	Size	Per-	Size	Per-	Size inches	Mineral	Per-	Size inches)	Mineral	Per- cent	Size (inches
1420	Quarts monsonite	. None.		Irregular.	One unit.	1/4	20		55		15		25											Biotite.	Trace.	1-1/2
1421	do.	do.		Lenticular	. do.	1/2	30		20		20		50											Kartite.	<1	
1422	do.			do.	do.	1	25		55		25		20											do.	Trace.	
1423	do.	None.		do.	Wall sone	2-3	30		50		70		20											do.	Trace.	
1424	do.			Irregular.		2-3	15		60		50		25											Magnetite.	<1 Trace.	1
1425	do.			Lenticular	Core.	3/4	30						100													
1426	do.			Irregular.		3/4	15		70		10		20 15 100											Martite.	<1	2015
1427	do.	Mone.		do.	One unit.	3/4	35		140		15		25									To all		Martite. Biotite.	<1 Trace.	
1426	Quarts monsonite and hornblende gneiss.	do.	Crosscutting.	Lenticular branching.	- Wall sone Core.	. 1	15		65		40		20				Trace.							Magnetite.	<1	
1429	Hornblende gneis		Not exposed.	Lenticular	Wall zone.	2	35		45		70		20				Trace.	3						Martite.	<1	
1430	Hornblende gneiss and quartz mon- sonite.	Mone.		Irregular.	Wall sone. Core.	1/4 24	70		5 20				15 80											Magnetite.	Trace.	1/16
1431	Quarts monsonite.			Lenticular	One unit.	1/16	70		10	1/4			20											Biotite.	Trace.	1/64
1432	do.	None.		Irregular.	do.	1/16	70		15		10	12	15											do.	Trace.	1/4
1433	do.	do.		Lenticular branching.	branch. North	2-1/2	¥5 85		40		30	12	15 15							4				Magnetite. Biotite. Magnetite.	<1 <1 <1	1/8
1434	do.			Lenticular	branch. One unit.	3/8	70		15	2			15							Act of						
1435	do.			do.	One unit.	1/4	65	2.39	20		5	3	15											do.	Trace.	1/16
1436	do.	None.		Lenticular- branching.	do.	1/2	55		30		20	5	15											Biotite. Magnetite.	<1 <1 Trace.	1/8
1437	do.			Irregular.	do.	1/2	65		20				15											Biotite.	Trace.	1/4
1438	Hornblende gneiss		Not exposed.	do.	do.	1-1/2	50		35		25	12	15											Magnetite.	Trace.	1/64
1439	do.		do.	do.	do.	1/8	65	2000	20		10	5	15											do.	Trace.	1/4
1440	do.		do.	do.	do.	1/8	70																	Biotite.	Trace.	1/32
1441	do.		do.	Lenticular.	do.	1/4	70		15		5	5	15				~1	1/64						Magnetite.		

Number and	Wall r	ock						Jerd Harr							Pe	matite			Sec.	V 44 41	STAIRS					
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture											Mi	neral	er.							
(Pl. II)	Tornacton		Wall Fock		structure	Inches/	Plagioc	lase	Perti	hite	Grapi	hic ite	Qua	rts	Muse	covite	Gar	net	Tours	aline	Lithium	mineral	•	Other	mineral	•
							Per-	Size nches)	Per-	Size nches	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size inches	Mineral	Per-	Size inches)	Kineral	Per- cent	Size (inches
1442	Hornblende gnei	•	Not exposed.	Lenticular- branching.	One unit.	3/4	55		30		20	10	15				Trace	. 1/8						Magnetite.	Trace.	1/8
1443	do.		do.	Lenticular	. do.	1/8	65		20		5		15						BASE!					do.	Trace.	1/16
1444	Quarts monsonite			Lenticular branching.	-Wall zone	12	25		55 10				20											do. Biotite.	Trace. Trace.	1 1/4 3
1445	Hornblende gneis		Not exposed.	Irregular.	One unit	1/8	70		15	2			15				<1	1/64						Magnetite.		1/4
1446	do.		do.	Lenticular	. do.	3/8	55		30				15											Magnetite.	Trace.	1/8
1447	do.		do.	do.	do.	1/2	50		35		5	8	15				<1	1/128						Biotite.	Trace.	1/16
1448	do.		do.	Irregular.	do.	1/16	68		15		5		15											do.	2	1/2
1449	do.		do.	Lenticular	. do.	1/16	75		10	1/2			15				Trace	1/64						do.	Trace.	1/16
1450	Quarts monzonite			do.	do.	1/4	65		20				15											Magnetite.	Trace.	1/4
1451	Hornblende gneis	s. None.	Not exposed.	Irregular.	do.	1/8	75		10	1			15											Biotite. do. Magnetite.	Trace. Trace.	1/4 1/16 3/8
1452	do.	do.	do.	Lenticular	. do.	1/8	45		40		10	8	15											do.	Trace.	1/8
1453	Quartz monzonite	. do.		Irregular.	do.	2	65		50				15											do.	Trace.	1/8
1454	do.			Lenticular	. Wall sone	. 3	40		35 10		25		2 5											do. Biotite.	Trace.	1/4 1/2
1455	Hornblende gneis	•	Not exposed.	Irregular.	One unit.	4	53		30		20	g	15											Biotite. Magnetite.	2 Trace.	1/16 1/32
1456	Quartz monzonite	. None.		Lenticular	. Wall zone	. 1/4	74		10				15											Biotite.	1	1/8
					Core.	2	25		55	5	TO A		20		Trace.									Magnetite.	Trace.	1/8
1457	do.	do.		do.	One unit.	1/2	75		10	2			15											do. Biotite.	Trace.	1/8 1/16
1450	ao.	do.		do.	do.	1	60		20	5			50											Magnetite. Biotite.	Trace.	1/16 1/32
1459	Hornblende gneis		Not exposed.	Lenticular branching.	- do.	3	65		20	10	8		15											Magnetite.	Trace.	1/16
1460	Quartz monzonite			Lenticular	Wall zone. Core.	8	25		40	20	12		35 60										6389	Biotite. Magnetite. Biotite.	Trace. Trace.	3/4 3/4 3
1461	Hornblende gneis	•	Not exposed.	Irregular.	One unit.	1-1/2	60		25		15	8	15											Biotite.	Trace.	1/4
1462	do.		do.	do.	do.	2	55		30		20	12	15											Magnetite. Biotite.	Trace.	1/16
1463	do.		do.	do.	do.	3	50	later !	35		25	12	15			-4							F	Magnetite.	Trace.	1/16

Number and	Wall r													- J 01 0	Pe	gmatite										
name of peg- matite	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture inches								46		Children .	M	neral	ogy							
matite (Pl. II)							Plagio	clase	Pert	hite	Grap		Que	erts	Musc	covite	Gaz	net	Tour	maline	Lithium	minera	ds	Other	mineral	
							Per-	Size (inches	Per-	Size	Per-	Size	Per-	Size	Per-	Size (inch	Per-	Size	Per-	Size	Mineral	Per- cent	Size	Mineral	Per- cent	Size (inch
1464	Hornblende gneis		Not exposed.	Lenticular	. One unit	. 4	30		55		45	12	15											Biotite. Magnetite.	Trace.	1/16
1465	do.		do.	Oval.	do.	2	55		25				20											do. Biotite.	Trace.	1/8
1466	do.		do.	Lenticular	. do.	3-1/2	35		50		40	8	15											Magnetite.	Trace.	1/4
1467	do.		do.	Irregular.	do.	5	65		20		10	12	15											Magnetite. Biotite.	Trace.	1/32
1468	do.		do.	do.	do.	2-1/2	42		37.5		30	12	15											do. Magnetite.	2 0.5	1/4
1469	do.		do.	Lenticula	r. do.	1-1/2	55		30		20	3	15											do.	Trace.	1/32
1470	do.		do.	Irregular.	do.	2	55		30		50	8	15		Trace	. 1/8								do. Biotite.	Trace.	1/2
1471	do.		do.	Oval.	do.	1	65		50		10	5	15											do.	Trace.	1/3
1472	do.	District St	do.	Lenticular	. do.	1/2	70		15		5	5	15				Trace.	1/32						Magnetite.	Trace.	1/32
1473	do.		do.	Oval.	do.	1-1/2	45		40		30	8	15				Trace.	1/32						do. Biotite.	Trace.	1/32
1474	do.		do.	Irregular.	do.	2-1/2	55		30				15											do. Magnetite.	Trace.	1/16
1475	do.		do.	Lenticular	do.	1	42.5		40		30	6	15											Biotite. Magnetite.	2 0.5	1/4
1476	do.		do.	do.	do.	5	45		40				15		Trace.	1/32	Trace	. 3/16	•					do.	Trace.	1/4
1477	do.	None.	Crosscutting.	do.	do.	2-1/2	45		40		30	8	15				Trace	. 1/16						do.	Trace.	1/4
1478	do.	do.	do.	Lenticular branching.	do.	1	50		30	3			50				Trace.	1/38						Biotite. Magnetite.	Trace.	1/8
1479	do.		Not exposed.	Lenticular	do.	1	60		25		15	8	15				Trace.	1/6								
1480	do.		do.	Lenticular- branching.	do.	3/4	70		15	4			15				Trace.	1/64						Biotite.	Trace.	1/8
1481	do.		do.	Lenticular.	do.	1/2	60		25		2		15				Trace.	1/64						do.	Trace.	1/4
1482	do.		do.	do.	do.	3/4	60		25	3			15	1			Trace.	1/64	•					Magnetite.	Trace.	1/61
1483	do.		do.	do.	do.	1	49.5		30	1650	5		20				0.5	1/38	4					do.	Trace.	1/8
1484	do.		dn.	do.	do.	2	63.5		20		10	8	15											Biotite. Magnetite. Biotite.	1 0.5	1/8 1/16 1/2
1485	do.		do.	do.	do.	1	49.5		30		20	12	15											do. Magnetite.	5 0.5	1/4
1480	do.		do.	do.	do.	2-1/2	36		45		35	12	15		1									Biotite. Magnetite.	3	1/4

hunber and	Wall r							al Carte	-							matite		11/4			Y LEWIS CO.				
ame of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	Texture inches										1		neral							
(F1. II)							Plagic	oclase	Pert	hite	Gran		Qu	arts	Musc	ovite	Gar	net	Tours	maline	Lithium	minerals	Other	mineral	
							Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Sise	Per-	Size	Mineral	Per- Size		Per-	Size (inches
1487	Hornblende gneis		Not exposed.	Lenticular	. One unit.	2	147		35		25	8	15										Biotite.	3	1/16
1488	do.		do.	do.	do.	1	65		20		10	8	14					T dell					do. Magnetite	1 Trace.	1/10
1459	do.		do.	do.	Wall sone. Core.	2-1/2	3674		35 80				20				Trace.	1/4					Biotite.	1	1/4
1490	do.		do.	Oval.	One unit.	1	60		25		20	5	15				Trace.	1/64					Biotite.	Trace.	1/8
1491	do.		do.	Lenticular	. do.	3/4	140		35		10		25						1				Magnetite		
1492	do.		do.	do.	do.	1-1/2	65		15		5		80		Trace.										
1493	do.		do.	Lenticular branching.	- do.	3	45		35	4	25	8	80				Trace.	1/16					Biotite.	<1 Trace.	2 1/32
1494	do.		do.	Lenticular	. do.	3	48.5		30		20	10	20		1	1/4	Trace.						do.	0.5	1/2
1495	do.		do.	Irregular.	Hanging wall layer. Footwall	4	25		45	8	5		30		Trace.	1/4	Trace.	1/16					do.	Trace.	1/32
					layer.	2	55		30		10	8	15				Trace.	1/16					do. Biotite.	Trace.	1/32
1496	do.		do.	Lenticular branching.	- One unit.	1/2	35		45		10		20		Trace.		Trace.								
1497	40.		do.	Lenticular	. Wall sone Intermediat sone. Core.	. 3/4	55		50		3	10	80 60 100		5 20	3									
1498	do.		do.	Lenticular branching.	- Wall some	. 1	55		30		20		15										Magnetite.	Trace.	
1499	do.		do.	Lenticular	Wall sone Core.	1/2	45		20 50				35 50		Trace										
1500	do.		do.	Lenticular branching.	One unit.	3/4	50		30				20		Trace.								Magnetite.	Trace.	
1501	do.		do.	Irregular.	Hanging wall layer Footwall	3	30		60		20		10										Magnetite.	Trace.	
					layer.	1/4	50		50				30										do.	Trace.	
1502	do.		do.	Lenticular	One unit.	2	55		25		15		20	2	Trace.								do.	Trace.	210
1503	do.		do.	do.	do.	4	35	1975	50		40		15		Trace.	1/4	Trace.								104
1504	do.		do.	do.	do.	2	10		60		60		30		frace.								Magnetite.	Trace.	
1505	do.		do.	do.	do.	1	50		30		40		20		Trace.								do.	Trace.	
1506	do.		do.	do.	Hanging wall layer Footwall	3	20		50				30										do.	Trace.	
					layer.	1/2	50		30				20										do.	Trace.	

Pegmatite

Trace. 1/64

<1 1/16

Wall rock

1531

1532

1533

do.

do.

do.

	Wall r	ock													Pe	gmatit	•			118 4		200				
Mumber and name of peg- matite	Type and formation	Alteration	Relation to	Shape	Internal	Texture								1111			M	neral	EY							
(P1. II)	TOPMECTOR		WALL FOCK		structure	Inches/		oclase	Pert	hite	Grap		Qua	arts	Muse	covite	Gaz	net	Tours	naline	Lithium	minera	als	Other	mineral	•
							Percent	Size						Size (inches			Per-			Size) Mineral	Per-	Size inches	Mineral	Per- cent	Sis (inche
1507	Hernblende gneis	•.	Not exposed.	Lenticular	.Wall sone,	1-1/2	55		30		20	g	15 60				Trace							Biotite.	Trace.	1/1
1508	do.		do.	Irregular.	Hanging wall layer. Footwall layer.	5	20		50		50		30		Trace		Trace							Magnetite.	∠1	
1509	do.		do.	do.	One unit.	2	39		4g				15		2	100	Trace									The same
1510	do.		do.	Lenticular	. do.	1/2	58		25				15		2	3							100			
1511	do.		do.	do.	do.	2	45		35	1	25		20		<1	1/4	Trace.	1/32								
1512	do.		do.	do.	do.	1	65		20		15	10	15		<1	2										
1513	do.		do.	do.	do.	1-1/2	55		25		20	5	20		<1	1/4	Trace.			9				Biotite.	Trace.	1/1
1514	do.		do.	do.	do.	3	40		45	10	35	5#	15				Trace.	1/16						Magnetite.	Trace.	1
1515	do.		do.	do.	do.	1	35		40		30	12	25	3336	Trace.	1/8	<1							do.	Trace.	1/
1516	do.		do.	Oval.	Wall some. Gore.	3/4	55		30	5	20	8	15 100				Trace.							do.	Trace.	
1517	do.		do.	Lenticular	One unit.	4			10				90				L Very							Monasite.	l crysta	1. 1/2
1518	do.		do.	Irregular.	Wall sone. Core.	3/4	55		30				15 100											Monasite.	l crysta	1. 3/1
1519	do.		do.	Lenticular	. One unit.	2	60		20				20		<1		Trace.						3			
1520	do.		do.	do.	do.	3/4	55		25		15	3	20		<1	1								Biotite.	<1	1/1
1521	do.		do.	do.	do.	1-1/2	45		35	5	5	5	20				Trace.									
1522	do.		do.	do.	do.	1	40		25		15		24							(Pract)				Biotite.	1	
1523	do.		do.	Irregular.	Wall zone. Core.	1-1/2	50		30		20	12	20											Magnetite. Biotite.	Trace.	1 2
1524	do.		do.	Lenticular	One unit.	3/4	65		20		15		15				<1									
1525	do.		do.	do.	do.	1/4	60		20				20		<1		Trace.	61-014						Biotite.	Trace.	
1526	do.		do.	do.	do.	3/4	65		50	3	5	8	15				<1	1/32	To see					do.	Trace.	1/:
1527	do.		do.	Lenticular branching.	do.	3/4	65		15		5	8	20				Trace.	1/64						do. Magnetite.	Trace.	1/1
1528	do.		do.	Irregular.	do.	1-1/2	50		35		25		15				Trace.	1/64					-	Biotite.	Trace.	1/1
1529	do.		do.	do.	do.	1-1/2	50		30		20	8	20				Trace.	1/64						do.	Trace.	1/8
1530	do.		do.	Lenticular	. do.	1/2	65		20		71-1	130	15				Trace.	1/64	-11				10.00	do.	Trace.	1/1

15

15

5 15

55

49

70

35

15

25

1-1/2

3/8

1/8

do.

do.

do.

do.

do.

do.

do.

do.

1/8

Trace.

Magnetite. Trace. Biotite, Trace.

	Wall r	ock					Size in			NA.					Peg	mat1t										
Number and name of peg-	Type and	Alteration	Relation to	Shape	Internal	Texture											M	neral	ey							
(Pl.II)	formation		wall rock		structure	inches	Plagi	oclase	Pert	hite	Grap		Quar	rts	Musc	ovite	Gaz	rnet	Tour	maline	Lithiu	n miner	ale	Other	mineral	•
							Per-	Size	Per-	Size	Per-	Size	Per-	Size inches)	Per-	Size	Per-	Size	Per-	Size	Mineral	Percent	Size	Mineral	Per-	Size (inches)
1534 The Trio Beryl Knob	Hornblende gneis		Not exposed.	Irregular.	Wall zone. Intermediate zone. Core. Intermediate zone.	3 36	50 20 50		35 50 20		30	10	15 10 80		20	2	<1							Biotite. Magnetite		1/8 1/4
1535	do.		do.	Lenticular branching.	Core Wall sone	10	60		50 25	5	Trace.		30 15		10	1/2	1140	1,-						Biotite. Magnetite. Magnetite.		1/4 1/8 1/8
1536	do.		do.	Lenticular	11 20 - 1		65		20	5	2		95		•	1/4	<1	1/16						Biotite.	Trace.	1/8
1537	do.		do.	Lenticular branching.	Carried States		50		35 10		5		15 89		1	1/2	<1							3101116	Irace.	1/6
1538	do.		do.	do.	One unit.	1	53	a lead	20				25		2											
1539	Covered by alluvium.		do.	Irregular.	Hanging wall zone. Footwall layer.	3 1/64	20		65	18	40	12	15		<1 race.	1/4	<1	1/16								
					Cora.	-4-	-		40				15 60		race.	1/0										
1540	Hornblende gneis		Crosscutting	do.	Core.	2 15	45		33 30	18	25	18	20 60		2	1/4								Gahnite.	۷۱	1-1/2
1541	do.	None.	do.	do.	One unit.	3/4	45		30	5			20		5	1/8	Trace.									
1542	do.	do.	do.	do.	Core.	1/16-1/1 8-12	83		2 55				15 37		1<1		Trace.									
1543	do.	do.	10.	Lenticular	Core.	6-8	65		15 35		Trace.		20 60		1<1		Trace.									
1544	do.		Not exposed.		Core.	3-4 3-4	29		50 55		60		20 37		1		Trace.									
1545	do.		do.	Lenticular	One unit.		64		15		10		20		1		<1									
1546	do.		do.	do.	do.	1/2	35		55				10		<1		Trace.									
1547	do.		do.	do.	do.	4	20		35				45													
1549	do.		do.	do.		1/2-1	30		49		45		20		1		Trace.									
1550			do.	Irregular.	do.	1/4-1/	59		20		10		20		1 <1		<1									
1551	dc.		do.	do.	North part South part. Core.	. 8	25 50		63 30 65	4 12			35 10 15 25		2 5	1/8	<1 <1 <1									
1552	do.		do.	Lenticular	Wall zone Core.	1/2	75		5	1			20				Trace.									
1553	do.		do.	Irregular.	One unit.	1/4	69.5	1/4	5	1/2			20		5	1/8	0.5	1/32								
1554	do.		do.	Ienticular	do.	1/2	抻		35				20	693	1		Trace.						Part of			

Mumber and	Wall ro	ock													Peg	matite	•	A STATE OF THE PARTY OF				Althor	Li-pits.			
name of peg-	Type and formation	Alteration	Relation to	Shape	Internal	Texture											M	neral	DEN.							
(P1.II)	Tornacton		wall rock		structure		Plagio	clase	Pert	hite	Grap	hic ite	Quar	ts	Musc	ovite	Gar	net	Tour	naline	Lithium	miner	als.	Other	mineral	•
							Per-	Size inches	Per-	Size inches	Per-	Size I	ent(1	Size	Per-	Size	Per-	Size	Per-	Size inches	Mineral	Per- cent	Size (inches)	Mineral	Per- cent	Size
1555	Hornblende gneis		Not exposed.	Irregular.	One unit.	1-2	20		64	75			15		1		<1	/								
1556	do.		do.	do.	do.	1/2-3/4	35		1414		15		20		1	4	Trace.							Biotite.	Trace.	
1557	do.	None.		do.	Wall zone. Core.	1/2	30		8		10		20 91		1 <1		Trace.							Columbite-	1 cryst	1. 1/2
1558	do.		do.	Lenticular	. Wall zone	.2-3	1 <1		78 20		90		20		1 <1									Columbite-	1 cryst	al. 1/h
1559	do.		do.	Lenticular branching.	- One unit.	2-3	10		75		75		15		<1		Trace.									
1560 Flood No. 1 and No. 2	do.		do.	Irregular.	do. Fracture filling.	1/2-1 8	30		149		5		20		1		Trace.							Beryl. 1 c	rystal.	2
1561	do.		do.	Lenticular	. One unit.	1	20		55		5		20		5		Trace.									
1562	do.	1	do.	Irregular.	do.	1	10		66		5		20		4	i i	Trace.									
1563	do.	None.	Crosscutting.	Lenticular	. Wall zone Core.	.1/2-3/1 6-8	50		35 50				15		۷1		۷1									
1564	do.		Not exposed.	do.	Wall sone. Core.	1/4	55 5		30 70				15 25		<1 <1		Trace.							Biotite. Magnetite.	Trace.	
1565	do.		do.	do.	One unit.	1-2	15		65		35		50		<1		Trace.									
1566	do.		do.	Irregular.	do. Fracture filling.	1/4-1/2	49		30		5		20		1		<1									
1567	do.		do.	do.	One unit.	3/4-1	15		64		20		20		1		Trace.							Martite. Biotite.	Trace.	
1568	do.		do.	do.	Wall zone. Core.	3-4 8	<1 2		80		85		20 98		<1 <1											
1569	do.		do.	do.	One unit.	1/4-1/2	15		63		10		20		2											
1570	do.		do.	do.	Wall zone. Core.	1/4-1/2	20		59		5		20 93		1 2		<1									
1571	do.		do.	Lenticular	One unit.	1/8	72		7				20		1		<1									
1572	do.		do.	do.	do.	1/2	69		3	2			20		7	1/8	1	1/32					200			
1573	do.		do.	Lenticular- branching.	Core.	1/2-3/4	10 8		60				30 75		<1									Beryl. 1	prystal.	1
157h Bucky Goldie South Slope	Hornblende gneiss and quartzite.	None.	Crosscutting.		Wall zone raphic gra ite inter- mediate zo	5	60		20 77				16 20		4		Trace.							Biotite.	<1	6
					Quartz- albite Intermediat	1/2	140		3	1-4			56		1	1/4-3/	4 <1									

Quarts

Graphic

grani te

Pegmatite

Muscovite

Mineralogy

Tourmaline

Lithium minerals

Other minerals

470 8 FS9

Garnet

Wall rock

Alteration

Relation to

wall rock

Not expessed.

Oval.

do.

1/2 40 40

10

2

<1

Trace

Shape

Internal

Texture

Plagioclase

Perthite

structure (inches)

Type and

formation

Mumber and name of peg-

(P1.II)

South Slope-

Continued.

1575

1576

1577

1578

1579

1580

1581

1582

1583

1584

1585

1586

1587

1588

1589

1590

1591

1592

1593

do.

Windy Knob

Camp Robber

matite

Mumber and	Wall ro	ock						6.724			The The			Peg	matite	•		Say Ye							
name of peg-		Alteration	Relation to	Shape		Texture										Mi	neral	NEA.							
(Pl. II)	formation		wall rock		structure	(inches)	Plagiocle	se P	erthite		phic nite	Quar	rtz	Musc	ovite	Gar	net	Tour	naline	Lithium	miner	als	Other	mineral	8
							Per- Si cent(inc	ze Pe	r- Siz	Per-	Size	Per-	Size nches)	Per-	Size (inche	Per-	Size	Por-	Size (inches	Mineral	Per- cent	Size	Mineral	Per-	Size (inches
1594	Hornblende gnei	6.	Not exposed.	Lenticular	One unit.	1/2-3/4	35	1	45	15		20		Trace		Trace.					100				
1595	do.			Irregular	Wall zone Core.	. 1	25		55	65		20		<1		Trace							Martite.	Trace.	
1596	do		Not exposed.	Lenticular	One unit.	8	25		54	30		50		1		Trace							Martite.	Trace.	
1597	do.		do.	do.	do.	1/8	65		15			50				<1									
1598	do.	None.		do.	do.	1/8	25		55			20		Trace		<1						177731			
1599	do.	do.	Crosscutting	. do.	do.	1/16	70	,	15			15				<1									
1600	do.		Not exposed.	Irregular,	do.	1-1/2	45	1	40	30		15		Part La									Magnetita,	Trace.	1/32
1601	do.		do.	Lenticular	. do.	1/4	75	1	10	5	8	15									114(1)		Magnetite,	Trace.	1/8
1602	do.		do.	Lenticular- branching.	do.	1/8	70	,	15	5	3	15									0,10		do.	Trace.	1/16
1603	do.		do.	Lenticular.	do.	1/8	58	-	25			15				2	1/32								
1604	do.		do.	do.	do.	1/8	75	1	10			15				< 1	1/16								
1605	do.	None.	Crosscutting	do.	do.	1/8	55	3	50	5		15				<1	1/64					*			
1606	do.		Not exposed.	do.	do.	1/8	50	3	30			20				<1	1/16					100			
1607	to.		do.	do.	do.	1/2	35	5	50			15				<1	1/32						4-17		
1608	do.		do.	do.	do.	1/8	75	1	10			15				Trace.	1/32								
1609	do.		do.	do.	do.	3/4	15	2	25			60		Trace.		<1							Biotite.	Trace.	
1610	do.		do.	Irregular.	Wall zone. Core pods.		35	5	55	50		20		<1		<1							do.	Trace.	
1611	do.		do.	Lenticular	.Wall zone. Core.	3-4	20		3	40		30 97		Frace.											
1612	do.		do.	Lentic- ular branching.	Wall zone. Core.	2-3	20	5	55			2 5 99		<1									Martite.	Trace.	
1613	Hornblende gneis and quartzite.		do.	do.	One unit.	3-4	30	4	15	30		25		Frace.		Trace.					11.5				
1614	do.		do.	Lenticular	do.	1-2	50	3	10	5		20		Frace.		Trace.									
615	Quartzite.		do.	do.	do.	1-2	40	14	10	5	1-36	20		<1		Trace.						1995			
616	Hornblende gneis		do.	do.	do.	1/4	50	2	25			25			1	<1							Magnetite.	<1	
1617	do.		do.	Irregular.	do.	3/4	20	6	io	15		20				Trace.							Biotite. Magnetite.	Trace.	
618	Quartz-biotite- schist.	None.	Crosscutting	do.	do.	2-3	20	5	55	55		25				Trace.							Martite.	<1	
1619	Hornblende gneiss and quartz-bioti schist.	te	Not exposed.	do.	do.	4-6	15	6	io	55		25											do. Biotite.	Trace.	

	Wall ro	ock							101/20	24.60					Peg	gmatite										
humber and name of peg-	Type and formation	Alteration	Relation to	Shape		Texture							niv di				M	neral	ey .							
(Pl. II)	Iormation		wall rock		structure(inches)	Plagio	clase	Perth	ite	Graph	hic ite	Quar	rts	Musc	covite	Gar	net	Tours	naline	Lithium	mineral	•	Other	mineral	
							Per-	Size (inche	Per-	Size Inches)	Per-	Size inches)	Per-	Size inches	Per-	Size	Per-	Size	Per	Size	Mineral		Size inches)	Mineral	Per-	Size
1620	Covered.		Not exposed.	Lenticular	. One unit.	1/4-1/2	50		30				20				<1							Biotite.	Trace.	
1621	Hornblende gneis		do.	do.	do.	1/4-1/2	45		35				20				<1	Up to						Biotite.	<1	
1622	do.		do.	do.	do.	1/4-1/2	35		45				20				Trace.		Trace					do.	<1	
1623	do.		do.	Lenticular branching.	- do.	1	15		50				35				∠1							do.	Tracel	
1624	do.		do.	Irregular.	do.	2	1		33				65		1		<1							do.	Trace.	
1625	do.		do.	Lenticular	. do.	1-2	15		60				25				<1									
1626	do.		Crosscutting.	do.	do.	3/4	30		45				25		<1		<1							Biotite.	Trace.	
1627	do.	None.		Irregular.	do.	1-1/2	155		7				35		3				Trace.		Lepidolite.	Trace.				
1628	do.		Not exposed.	Lenticular	. do.	1/4-1/2	35		45				20		Trace.	200	<1							2.0		
1629	do.		do.	Lenticular branching.	-Wall zone.	1/2	40		40				20		Trace.									Biotite. Magnetite.	Trace.	
1630	do.	None.		do.	Core. Wall zone.	1/4	20		10 55				90		<1 <1		<1							Magnetite.		
					Core.	3-6	1		5				94				Trace.								1 cryst	al.
1631	Tonalite.	do.		do.	Wall zone. Core.	1/4 3-4	25 <1		50 7				25 93				<1									
1632	do.	do.		Irregular.	Wall zone.	1/8	30		40				30				Trace.							Biotite.	Trace.	
					Core.	3			5		1		95					ي ساود						Martite.	Trace.	
1633	io.	40.		do.	One unit.	1/2-3/4	25		50				25				Trace.							Biotite.	Trace.	
1634	Hornblende gneis		Not exposed.	Lenticular	.Wall zone. Core.	3-4	15		65 Trace.		40		20		Trace.											
1635	do.		do.	Oval.	One unit.	1/4	25		50				25				Trace.							Biotite. Martite.	Trace.	
1636	do.		do.	Lenticular	. do.	1/8	35		717				30											Biotite. Martite.	1 <1	
1637	do.		do.	do.	do.	1/2	15		60		26 70		25											Biotite.	<1 <1	
1638	Covered.		do.	Irregular.	Wall zone.	1	20		55		60		25											Biotite.	Trace.	
639	Bornblende gneise	None.		Lenticular	One unit.	1/2	40		35				25		<1		<1							Biotite.	Trace.	
1640	do.	do.		do.	do.	1	40		30				30		<1		<1									
1641	do.	do.		do.	do.	3/4	15		65				20		<1		<1									
1642	do.		Not exposed.	Irregular.	Wall zone. Core.	1/2-3/4	35		40		5		25 100				<1							Biotite.	<1	Up t
1643	Hornblende gneiss		do.	do.	One unit.	1/2	20		55				25	103	Trace.		<1	13.3		C. A. P.			Ser.	Biotite.	Trace.	

	Wall r	ock													Peg	matite				RES.	Mary United					
umber and ame of peg-	Type and	Alteration	Relation to	Shape		Texture											Ms	neral	ову				192			
atite (Pl. II)	formation		wall rock		structure	(inches)	Plagio	clase	Pert	hite	Grap		Quar	rtz	Musc	ovite	Gar	net	Tour	aline	Lithium	minera	18	Other	mineral	•
							Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Mineral	Per- cent (Size inches)	Mineral	Per- cent	Size (inches
1644	Hornblende gnei	18,	Not exposed.	Lenticula	. One unit	1/4	20		55				25		Trace		<1							Biotite.	<1	
1645	do.		do.	do.	do.	1/8-1/4	35		35				30				<1							do.	<1	
1646	do.		Crosscutting	do.	do.	1-2	15		50				35				<1							do.	<1	
1647	do.		do.	Lenticular- branching.	do.	1/4-1/2	35		40				25		Trace		<1							do. Magnetite.	Trace.	
1648	Fornblende gneis	a None.		do.	do.	1/4-1/2	30		35				35		Trace		<1							Biotite.	Trace.	
1649	Hornblende gnei	٠.	Not exposed.	Lenticular,	do.	1/4-1/2	15		55				30				<1							do.	Trace.	
1650	do.		do.	do.	do.	1/2	40		40				20				Trace.							do.	Trace.	
1651	do.		do.	Irregular.	do.	1/4	35		40				25				Trace							do.	Trace.	
1652	do.		do.	Lenticular- branching.	do.	1/2	30		50		5		20				frace.							do.	Trace.	
1653	do.		do.	Lenticular.	do.	2-3	20		65		10		15											Magnetite.	<1	
1654	do.		do.	do.	do.	<1/8	19		55				25	11-11-1										Biotite.	,	
1655	do.		do.	Lenticular- branching.	do.	1/8-1/4	40		39				20											do.	1	
1656	do.		do.	do.	Wall zone. Core.	6-8	25		55		50		20		<1									do.	<1	
1657	do.		do.	Irregular.	One unit.	1	20		55		30		25											Biotite. Martite.	<1 <1	
658	do.		do.	Oval.	do.	1-2	10		70		10		20											Biotite.	<1 <1	
659	do.	None.		Lenticular branching.	- Wall sone	.1-2	30		45		30		25											Biotite.	<1	40
				or anoming.	Core.	12			15				85											Martite.	<1	
1660	do.		Not exposed.	Irregular.	One unit.	2-3	15		60		75		25											Biotite. Martite.	. <1 <1	
1661	Quartz monzonite.	None.		do.	Wall zone.	4	10		70				20	E										Biotite.	<1	
					Core.	1							100											Magnetite.	<1	
.662	do.	do.		do.	One unit.	2-3	40		40		15		20											Biotite. Martite.	<1 <1	
663	do.	do.		do.	do.	1	50		25			2	55											Biotite.	Trace.	
664	do.			Lenticular	do.	1/4	30		50				20											do. Martite.	<1 <1	
	do.			do.	do.	1/8	20		55				25											Biotite.	<1 <1	
666	Hornblende gneiss	. None.		Irregular.	Wall zone.	1/2	20		50		15		30											Biotite. Martite.	7 77	

Number and	Wall r	ock			1 Pol (5 - 27)				2.14.32	STORY.			della		Pe	matite	1860		1100	0.1566.1		100				
name of peg- matite	Type and formation	Alteration	Relation to wall rock	Shape		Texture																				
(Pl. II)					structure	inches)	Plagioclase		Perthite		Grap		Quartz		Muscovite		Garnet		Tourmaline		Lithium minerals			Other minera		•
								Size inches				Size (inches	Per-	Size	Per-	Size	Per-	Size	Per-	Size (inches	Mineral	Per- cent(Size inches)	Mineral	Per- cent	Size (inches
1667	Quarts monzonite		1	Irregular.	One unit.	1/8	10		55				35											Bictite. Martite.	<1 Trace.	
1668	do.			Lenticular branching.	- do.	1/16	72		2				25											Biotite.	1	
1669	do.	None.		Irregular.	do.	3/4-1	2		78		5		20											do.	Trace.	
1670	do.			Lenticular	. do.	4-6	20		80		85	1 0.	20											do.	Trace.	
1671	Hornblende gneis		Not exposed.	do.	do.	1/8	55		20				25											Biotite. Martite.	<1 <1	
1672	do.		do.	do.	dc.	2-3	30		45		25		25											do.	Trace.	
1673	do.		do.	Irregular.	do.	1/8-1/4	50		25				25				Trace.							Magnetite.	Trace.	
1674	do.		do.	Lenticular branching.	- do.	1/4	30		140				30		<1		<1									
1675	do.	None.		Irregular.	One unit.	1/2	5		70				25				<1							Biotite.	Trace.	
1676	do.	do.	Crosscutting.	do.	do.	1/8-1/4	60		15				25				Trace.							do.	<1	
1677	do.		Not exposed.	do.	do.	1/8-1/4	35		40				25				Trace.							do.	<1 Trace.	
1678	do.		do.	Lenticular	do.	1/8	64		10				25											do. Biotite.	1 <1	
1679			do.	do.	do.	1/8-1/4	65		15				25											do.	Trace.	
1680	40.		do.	do.	do.	1/4	50		30				20											Magnetite.	Trace.	
1681	do.		do.	Irregular	do.	1/4	35		40				25											do. Biotite.	<1 <1	
1682	Quartz monzonite.	None.		do.	do.	3/4	20		55				25											Magnetite.	<1	
1683	Fornblende gneiss		Not exposed.	do.	do.	1/8	60		15				25											Biotite. Magnetite.	<1 Traco.	
1684	do.		40.	do.	do.	1/2-3/4	15		55				30		Trace.		<1							Biotite.	<1	
1685	do.		do.	Lenticular	do	1/4-1/2	3		61				35		<1		1									
1686	do.		do.	do.	do.	1/4-1/2	10		60				30				<1							Biotite.	<1	
1687	de.		do.	do.	do.	1/2	10		55				35				<1							do.	<1	
1688	do.		do.	do.	do.	1/4-1/2	7		63				30				<1									
1689	do.	10 10 10	do.	do.	do.	1/4-1/2	5		60				35				<1									
1690 Lazy Day	do.		do.	do.	do.	1/4	40		35				25		<1		<1							Biotite.	Trace.	
1691	40.		do.	do.	do.	/4-1/2	20		55	1 1 12			25				<1	3793						do.	Trace.	
1692	do.		do.	Irregular.	do.	1/4	12		63				25				<1					100		do.	<1	

Number and name of peg- matite (F1.II)	Wall rock		No.							T.					Pog	matite								A PROSPER		
	Type and formation .	Alteration	Relation to	Shape	Internal structure	Texture	Mineralogy																			
			wall rock			(inches)	Plagioclase		Perthite		Graphic granite		Quarts		Muscovite		Garnet		Tourmaline		Lithium minerals		ls.	Other miners		
							Percent	Size	Per-	Size inches	Per-	Size inches	Per-	Size	Per-	Size	Per-	Size	Per-	Size	Mineral		Size inches)	Mineral	Per-	Sis (inch
1693	Hornblende gneis and quarts mon- zonite.		Not exposed.	Lenticular branching.	- One unit.	1/4-1/2	45		35		15		50													
1694	Fornblende gneis		do.	Lenticular	. do.	1-1/2	10		70		65		20													
1695	Fornblende gneis and quartz mon- zonite.		do.	do.	do.	1	15		65		40	14	20						issi Here					Magnetite.	Trace.	
1696	do.	None.		Lenticular branching.	- West part	1-1/2	60 20		15 60	2	₹ 5		25		Trace.									Biotite.	Trace.	
1697	Fornblande gneis	١.	Not exposed.	do.	One unit.	1	10		70		70		20							Dr. 1						
1698	do.		do.	Lenticular	. do.	2	10		70		80		20											Magnetite.	Trace.	
1699	Quartz monzonite			do.	do.	1	15		65		70		20			State of			THE LINE					do.	Trace.	
1700	do.			do.	do.	2	5		75		85		20		4-1											
1701	Quartz monsonite and hornblands gnaiss.		Not exposed.	Lenticular branching.	- do.	1-2	10		70		70		20													
1702	do.		do.	Lenticular	do.	1	10		70		85		20											Magnetite.	Trace.	
1703	Hornblande gneis		do.	Lenticular branching.	- do.	2	15		65		70	Up to	20											Biotits.	Tracs.	
1704	do.		do.	Lenticular	. do.	5	50		55		50	Up to	25													
1705	do.		do.	do.	do.	1/2-1	45		35		25		20													
1706	Fornblende gneise and quartz mon- zonite.		do.	do.	do.	1-1/2	30		50		60		20											Magnetite.	Trace.	
1707	Hornblende gneis		do.	Lenticular branching.	- do.	1/4	60		25	Up to	1		15											Biotite. Magnetite.	Trace.	1/2
1708	do.		do.	Irregular.	do.	1/4-1/2	45		30	Up to	5		25											do.	Trace.	
1709	do.		do.	Lenticular	do.	1/2-1	50		30	Up to	10		20											Biotite. Magnetite.	Trace.	1/4
1710	do.		do.	do.	do.	1/4-1/2	60	1	20		10		20				Trace.							do.	Trace.	
1711	do.		do.	do.	do.	1	35		45		40		25											do.	Trace.	
1712	do.		do.	Irregular.	do.	1/2-1	40		40		20		20											Biotite. Magnetite.	Trace.	
1713	do.		do.	Lenticular branching.	do.	1/4	50		30				50		Trace.		Trace.							Biotite. Martite.	Trace.	
1714	do.		do.	do.	do.	1/2	145		35		5	6-10	20				Trace.							do.	Trace.	LE BE

1/4	
< 1/2	
1/2	
1/4-1/2	TC2
3/4	
1/8 1/32 1/32-1/4	
1/16 1/8 1/32	
1/8 1/32	
1/4 1/32	
1/8	

Number and name of peg- matite	Wall rock			Pagnatite																							
	Type and formation	Alteration	Relation to wall rock	Shape		Texture (inches)	hes)																				
(Pl. II)					structure		Plagioclase		Perthite		Gray	Graphic granite		Quarts		Muscovite		net	Tourmaline		Lithium minerals		ı.	Other miner		al.	
							Per-	Size	Per-	Size	Per-	Size (inches	Per-	Size	Per-	Size	Per-	Size inche:	Per-	Size	Mineral	Per-	Size (inches	Mineral	Per-	Size (inches	
1715	Hornblende gneie		Not exposed.	Lenticular	. One unit	1/4	60	1/16-	20	1/2-2			20				Trace.										
1716	do.		do.	do.	do.	1/4	60		20	1/2-1			20				Trace.										
1717	do.		do.	Lenticular branching.	- do.	1	35	1/5	45	1-3	20		20	1/8-1/										Magnetite.	<1	1/4	
1718	do.		do.	do.	do.	1	50		30	1/2-3	30		20											do.	<1	< 1/2	
1719	do.		do.	Irregular.	do.	1/4	60	1/8	20	1/2-			10	1/4-1													
1720	Hornblende gneis and quarts mon- sonite.	None.			Northeast and South- rest ends. Central part.	1	35		79		10		20				Trace.							Martite.	<1	1/2	
1721	Quarts monzonite.	do.		Lenticular				1/16-		1/2-3	30	1-3	20	1/5-1/	Trac	e.1/8	<1 Trace.	1/8						Magnetite. Biotite.	<1 Trace.	1/4-1/2	
1722	Hornblende gneiss and quarts mon- scnite.	do.		Lenticular branching.	do.	2	30		50		50		20				Frace.							do. Magnetite.	Trace.	3/4	
1723	do.			do.	do.	5	5		75		40		20				<1							Biotite.	Trace.		
1724	Hornblende gneise		Not exposed.	Lenticular	do.	3	30		50		30		20				Frace.	1/32						Magnetite.	Trace.	1/8	
1725	do.		do.	do.	do.	3/4	60		25		10	6	15												Trace.	1/32	
1726	do.		do.	do.	do.	1/4	50		30	3			20				<1	1/16						do.	<1	1/32-1	
1727	do.		do.	do.	do.	3/4	35		45		10	5	20				Trace.	1/16								-, ,,	
1728	do.		do.	Irregular.	do.	4	39		50		60	10	20											Biotite.	Trace.	1/16	
1729	Hornblende gneiss and quarts mon- sonite.		do.	Lenticular	do.	1/2	45		35	3			20												Trace.	1/8	
1730	do.		do.	do.	do.	3	30		55		40	6	15												Trace.	1/8 1/32	
1731	Hornblende gneiss		do.	do.	do.	1/4	49.5		30	3			20				0.5	1/16						Biotite.	Trace.	1/4 1/32	
1732	Fornblende gneiss and quarts mon- sonite.		do.	40.	do.	1/2	65	1/8	20		20	14	15												Trace.	1/32	
1733	do.		40.	Lenticular- branching.	do.	1/4	45		35		20	4	20											Magnetite.	Trace.	1/16	
1734	do.		do.	do.	do.	6	10		70		85		20		Trace									Biotite. Martite. Samarskite	<1 <1 (1) 3		
1735	Fornblende gneiss		do.	do.	do.	1/2	55	1/32	30	4			15	300	2654					193				Magnetite.		1/16	

Number and	Wall r	ock											200		Pog	matite								A MESSAGE		
ame of peg-	Type and	Alteration	Relation to	Shape	Internal structure(Texture											N	ineralo	EY							
natite (Pl. II)	formation		wall rock		structure	Inches)	Plagic	clase	Pert	hite	Gray		Qua	arts	Musc	covite	Gas	rnet	Tours	aline	Lithium	miner	de	Other	mineral	
							Per- cent(Size inches	Per-	Size inches	Per-	Size	Per-	Size inches	Per-	Size inches	Per-	Size	Per-(Size inches	Mineral	Per- cent	Size (inches	Mineral	Per- cent	Sise (inch
1736	Hornblende gneis and quarts mon- sonite.		Not exposed.	Lenticular	. One unit	2	30		50	ħ	30	6	20											Martite.	Trace.	1/8
1737	do.		do.	do.	do.	1	50		30	3	15	14	20				<1	1/64				1		do.	Trace.	1/32
1738	do.		do.	Lenticular branching.	- do.	2	40		34.5		40	8	25				0.5	1/8						Magnetite.	Trace.	1/32-2
1739	Quartz monsonite			Lenticular	. do.	1/2	30		50		50		20			8,11								do. Biotite.	Trace. Trace.	1/64
1740	Quarts monsonite and hornblends gneiss.		Not exposed.	do.	do.	2	50		35		50	8	15											do. Magnetite.	Trace.	
1741	Quarts monsonite			do.	do.	3	30		50		65	8	50											do.	Trace.	1/32
1742	do.			do.	do.	1-1/2	35		50		35	5	15													
1743	do.	None.		Irregular.	do.	2	40		45		30	5	15											Biotite.	Trace.	1/4
1744	do.			Lenticular	. do.	2	40		50		40	4	10											Magnetite. Biotite.	Trace.	1/8
1745	do.			do.	do.	1/2	60		25		50	5	15											Magnetite.	Trace.	1/16-1
1746	do.			do.	do.	3/4	50		30		30	5	20											do.	Trace.	1/16
1747	do.			do.	do.	5	25		65		70	8	15													
1748	do.			do.	do.	1	55		35	4	15	5	20				<1	1/64				1.56		Magnetite.	Trace.	1/16
1749	do.			do.	do.	2	30		55		40	8	15			601.81								do.	Trace.	1/16
1750	do.			do.	do.	1	35		45		20	8	20				Trace	1/16						do. Biotite.	Trace.	1/32
1751	do.			do.	do.	4	20		60		60	8	20											Magnetite.	Trace.	1/16
1752	do.			Lenticular branching.	- do.	1/2	55		30	2			15											do.	Thace.	1/4
1753	do.			Lenticular	. do.	4	25		50		50	10	25											do.	Trace.	1/32
1754	do.			do.	do.	1	55		35	14	15	5	50				<1	1/64						do.	Trace.	1/16
1755	do.			do.	do.	5	30		50		50	5	20											do.	Trace.	1/8
1756	do.			do.	do.	4	25		60		60	6	15											Riotite. Magnetite.	Trace.	1/8 3/16
1757	do.			Lenticular- branching.	do.	5	30		55		50	8	15				Trace	1/64						Biotite. Magnetite.	Trace.	1/32
1758	do.			Lenticular	do.	4	30		55		50	10	15				Trace.	1/32						do.	<1	1/32
1759	do.			Irregular.	do.	3	20	1	65		70	8	15			21/4						The same	B	do.	Trace.	1/8
1760	Quartz monzonite and hornblende gneiss.		Not exposed.	Lenticular	do.	1-1/2	50		65		60		15											do.	Trace.	1/4

	Number and	Wall r	ock									Stupe	The .			Peg	matite									PAPER.	138
Principal Prin	name of peg-	Type and	Alteration		Shape														tineral	ogy							
	(Pl. II)	Iormation		wall rock		structure	(inches)	Plagic	oclase	Pert	thite			Que	arts	Muso	ovite	Ga	arnet	Tous	maline	Lithium	n miners	ıle.	Other	mineral	•
								Percent	Size	Per-	Size	Per-	Size (inche	Per-	Size	Per-	Size	Per-	Sise	Per-	Size	Mineral	Per-	Size	Mineral		Size (inches)
1763	1761	Quartz monzonite				One unit.	3	45		140		40	8	15											Nagnetite.	Trace.	1/16
176	1762	do.			Lenticular	. do.	2	45		60		20	6	15				Trac	1/32								1/16 1/32
1766 do. d	1763	do.			do.	do.	2	25		6c		20	5	15		08/01/01		Trace	1/32			1.4					1/16
176	1764	do.			do.	do.	3	30		55		50	8	15													1/8
1766 do. do. do. 1 b5 b0 10 5 15 do. Trace. 1767 do. do. 30 30 65 66 22 15 do. Trace. 1768 do. Trace. 1766 do. Tra	1765	do.			do.	do.	2	30		55		40	5	15												Marie en	1000
176	1766	do.			do.	do.	1	45		40		10	5	15	61200												1/16
1768 do.	1767	do.			do.	10.	3	50		65		6c	12	15													1/16-1
1770	1768	do.			Lenticular branching.	- do.	3	15		65		60	8	20				•							1114		1/16
1771	1769	do.		, '	Lenticular	. do.	2	30		55		50	8	15					3.7			•			do.	Trace.	1/16
1772 do.	1770	do.				- do.	2	30		55		50	5	15											do.	Trace.	1/32
1772 do. Noss. do. do. 3 35 50 60 8 15	1771	do.			Irregular.	do.	1/2	40		35		5		25											Martite.	<1	
	1772	do.	None.		do.	do.	3	35		50		60	8	15										STATE	Magnetite.		1/32
1775 do. do. do. do. do. do. do. do.	1773	do.				- do.	1/2	40		40		10		20											Martite.		
1775 do. do. do. 1/4 55 30 3 10 4 15 15 16 do. 1776 do. do. do. 2 25 55 30 6 20 1776 do. 1776 do. do. do. 40. 40. 5 15 15 15 1776 do. 1776 do. 1776 do. do. do. 1 50 30 20 5 20 1779 do. 1776 do	1774	do.			Lenticular	do.	3	10		75		70		15						2					Magnetite.	Trace.	1/8
1776 do.	1775	do.			do.	do.	1/4	55		30	3	10	4	15											Biotite.	Trace.	1/16
1777	1776	do.			do.	do.	2	25		55		30	6	20						Tay					do.	Trace.	3/16
1778 do. do. do. do. 1 50 30 20 5 20 do. Trace. 1779 do. do. do. do. 1 30 55 20 65 70 8 15 do. Trace. 1780 do. Lenticular do. 1/2 25 60 3 15 do. Trace. 1782 do. do. do. 2 25 50 10 5 25 do. Trace. 1783 do. do. do. 3 40 40 80 8 20 do. Trace. 1784 do. do. do. 3 30 50 40 8 20 do. Trace. 1785 do. do. do. 1/4 70 15 3 15 do. Trace. 1786 do. do. do. 5 30 40 80 8 20 do. Trace. 1786 do. do. do. 5 30 40 80 8 20 do. Trace. 1786 do. do. do. 5 30 80 80 80 80 80 80 80 80 80 80 80 80 80	1777	do.		7-15	do.	do.	4	35		50		40	5	15											Magnetite.	1000	3/8
1779 do. do. do. do. do. do. do. do.	1778	do.			do.	do.	1	50		30		20	5	20											do.	100	1/32
1780 do. Lenticular do. 1 30 55 20 15 do. < 1	1779	do.			do.	do.	5	20		65		70	8	15											do.	Trace.	1/4
1782 do. do. do. 2 25 50 10 5 25 do. Trace. 1783 do. 1784 do. 1785 do. do. do. 1/4 70 15 3 15 do. Trace. 1785 do. 1786 do. do. do. do. 5 30 ho s 20 Biotite. Trace. Magnetite. Trace. Magnetite. Trace.	1780	do.				do.	1	30		55		20	1	15											do.	<1	
1782 do. do. do. 2 25 50 10 5 25 1783 do. 1784 do. do. do. do. 1/4 70 15 3 15 1785 do. 1786 do. do. do. do. 5 30 ho 5 20 1786 do. do. do. 5 30 ho 5 20 Biotite. Trace. Magnetite. Trace. Magnetite. Trace.	1781	10.			Lenticular.	do.	1/2	25		60	3			15											do.	Trace.	1/8
1763 do. do. do. do. do. do. do. 1/4 70 15 3 15 15 16 do. Trace. 1/765 do.	1782	do.			do.	do.	2	25	1	50		10	5	25		Egg 3							100			300	1/4
1784 do. do. do. 1/4 70 15 3 15 15 16 do. Trace. 3/ 1785 do. do. do. do. 5 30 ho s 20 Biotite. Trace. Magnetite. Trace. Trace.	1783	do.			do.	do.	3	40	100	40		40	8	20											do.	133 633	1/8
1785 do. do. do. 3 30 50 40 8 20 Biotite. Trace. Magnetite. Trace.	1784	do.			do.	do.	1/4	70		15	3		E YE	15													3/16-1/2
1786 do. do. do. 5 30 h0 70 10					do.	do.	3	30		50		40	8	20											Biotite.	Trace.	3/8 1/16
	1786	do.			do.	do.	5	30		40		30		30											do.		1/8-2

253

unber and		47444	2-1-44				PEC C	-	X 5 1/4	1005						matite						5757				
me of peg-	Type and formation	Alteration	Relation to	Shape	Internal structure	inches)										1		neralo	The state of							
P1.II)								clase		hite	Grap	ite	157	rts		ovite		rnet		aline	Lithium	minera	1.	Other	mineral	•
							Per- cent	Size	Per-	Size inches	Per-)cent(Size inches	Per-	Size inches	Per-	Size inches	Per-	Size inches	Per-	Size inches) Mineral	Per- cent(Size inches)	Mineral	Per- cent	Sise (inches
1787	Quarts monsonite			Lenticular branching.	One unit.	2-3	10		70		80		20											Martito.	~ 1	
1788	do.			Lanticular	. do.	1-2	30		50		45		20											Magnetite.	~ 1	
1789	do.	4		do.	do.	1-2	20		55		40		25											do.	~ 1	
1790	do.			Lenticular branching.	- do.	1/8-1/4	55		25				20				<1							Kartite.	Trace.	
1791	do.			Lenticular	. do.	1/8-1/4	70		10				20				<1									
1792	do.			do.	do.	1/2	45		35		20		20				Trace			, T.J.				Martite.	∠1	
1793	do.			do.	do.	3/4	35		40		2		25				<1							do.	Trace.	
1794	do.			Lenticular branching.	- do.	1/2	50		25		1		25				Trace							do.	<1	
1795	do.			Lenticular	. do.	3/4	35		45		Trace.		20				Trace							do.	Trace.	
1796	do.			Lenticular branching.	- do.	1/4	50		30		2		20				Trace							Magnetite.	Trace.	
1797	do.			Lenticular	. do.	1	15		65		60		20											do.	1	
1798	do.			Lenticular branching.	- do.	2	25		55		30		20				<1							Biotite. Nartite.	Trace.	
1799	Hornblende gneis	s. None.	Crosscutting.	Lenticular	. do.	1/2-3/4	30		40				30		Trace.											
1500	Quartz monzonite	. do.		Irregular.	do.	3/4	25		55		5		20											Biotite.	Trace.	
1801	Hornblende gneis	s. do.		do.	do.	1/8-1/4	63		8				27		2		Trace.							Martite.	<1	1/8-1-
1802	do.			Lenticular branching.	- do.	1/8-1/4	70		9				20		1		Trace.						-			
1803	do.			do.	do.	1	63		20				15		s											
				745																						
						2340																				
															4P - 1											
							46	1200	Bar.		100	Surgery !	NOTE OF	199		20, 34	Flee 3	Sec.	The state of			1200	8.39			

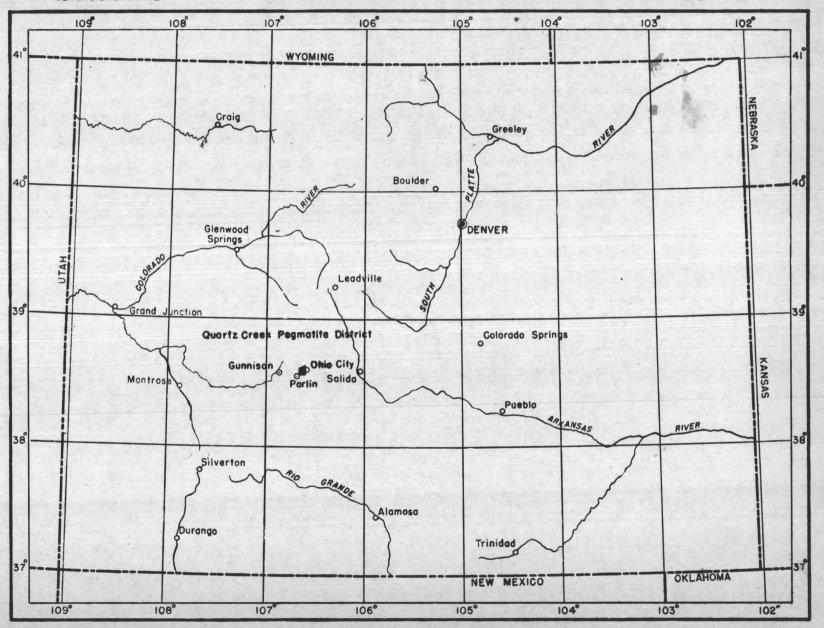


FIGURE I. INDEX MAR QUARTZ CREEK PEGMATITE DISTRICT, COLORADO

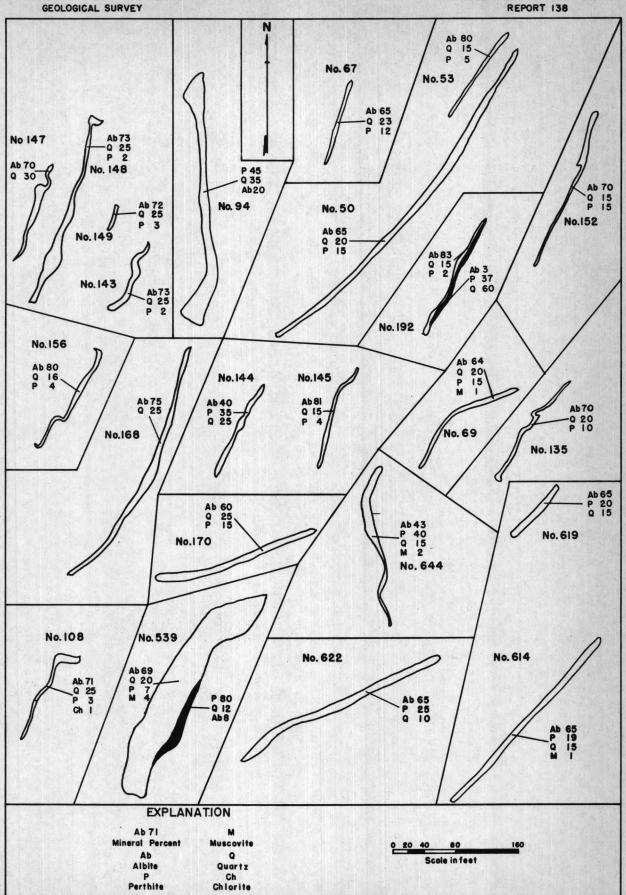


FIGURE 2. LENTICULAR PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

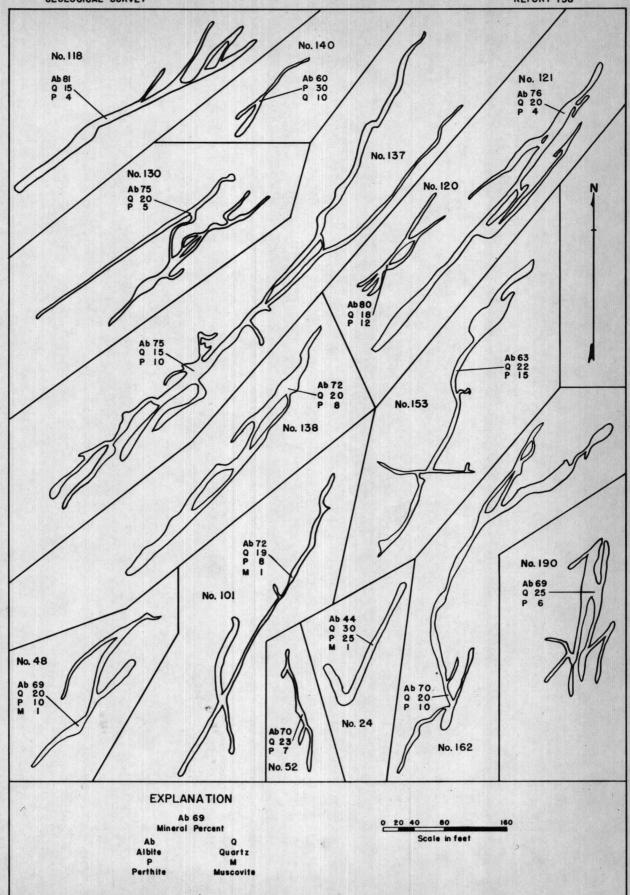


FIGURE 3. LENTICULAR AND BRANCHING PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

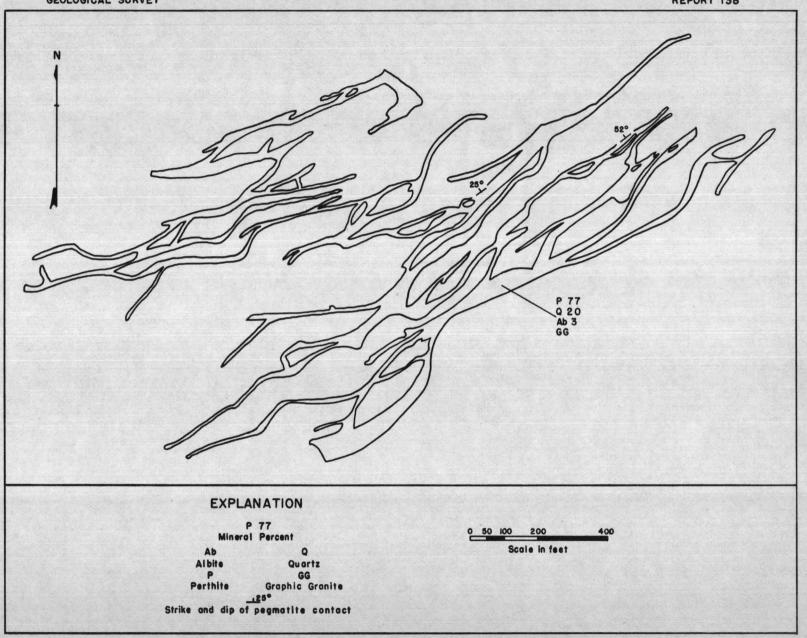


FIGURE 4. SHAPE OF PEGMATITE NO. 1294, QUARTZ CREEK PEGMATITE DISTRICT

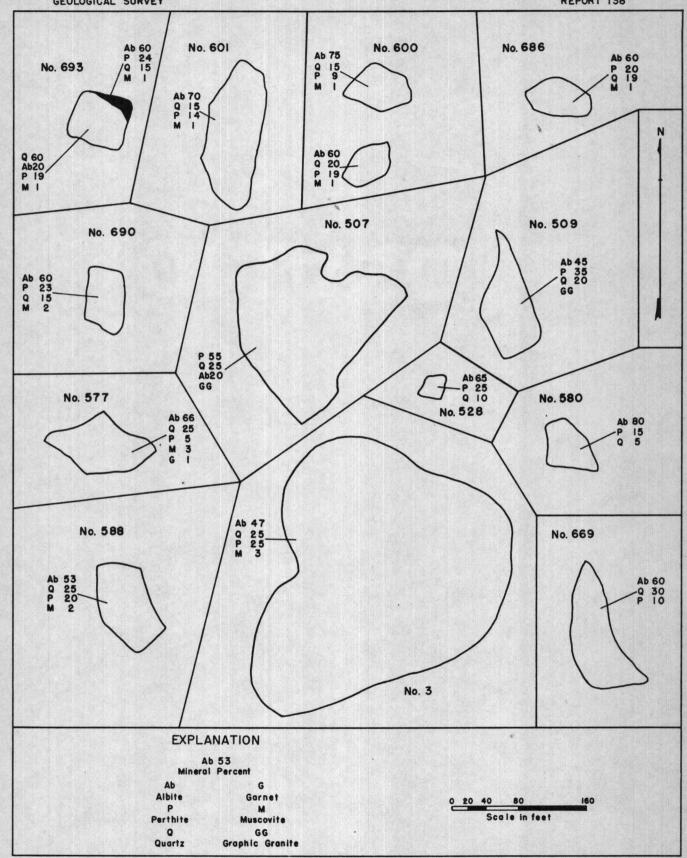


FIGURE 5. OVAL PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

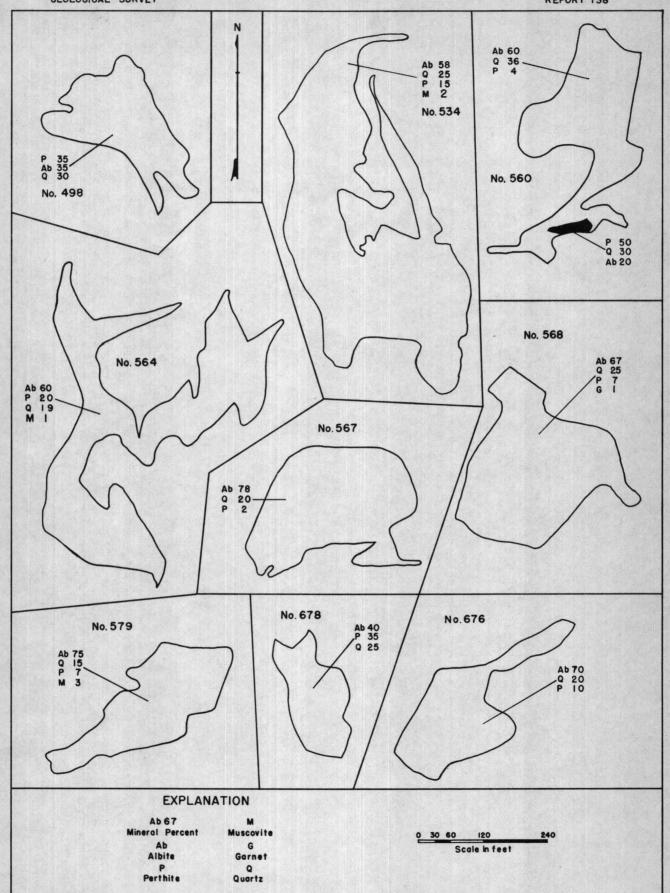
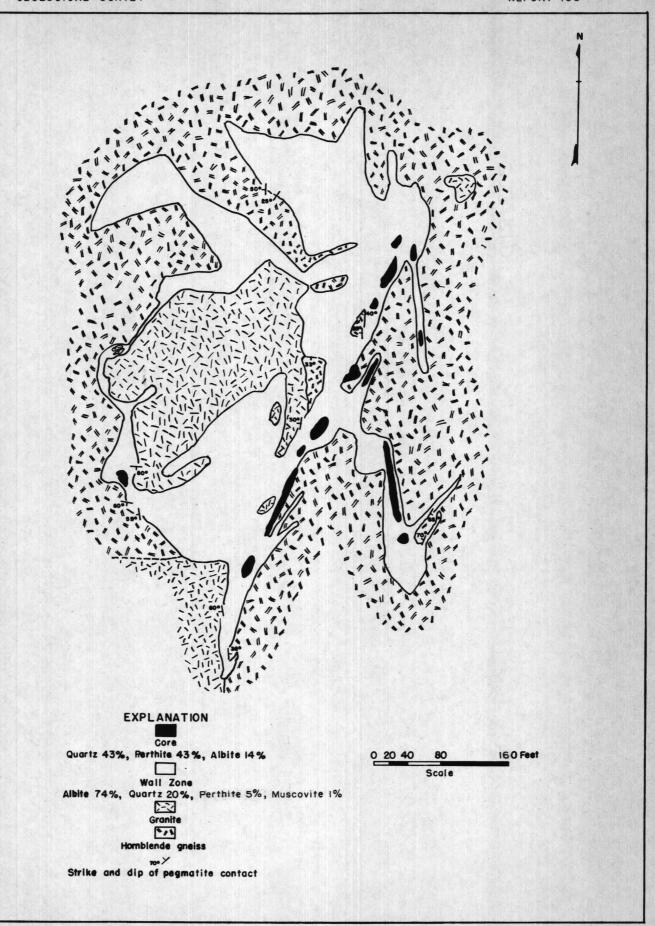


FIGURE 6. IRREGULAR PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

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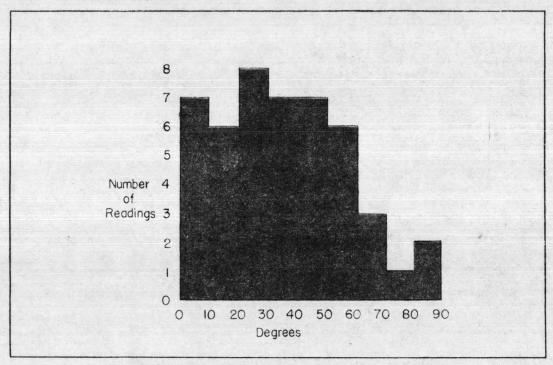


FIGURE 8. ANGLE BETWEEN PEGMATITE CONTACTS
AND FOLIATION OF COUNTRY ROCK

FIGURE 10. ZONED PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

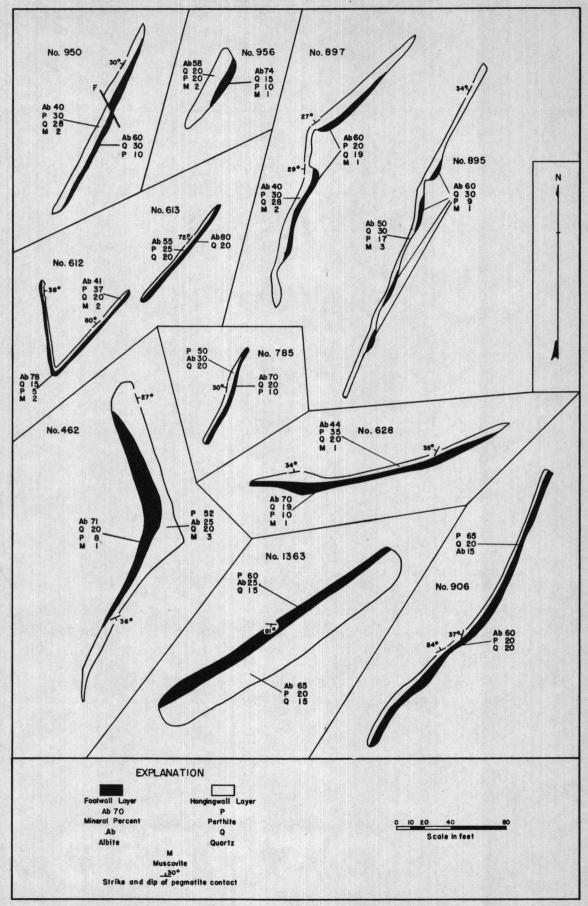


FIGURE 11. LAYERED PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

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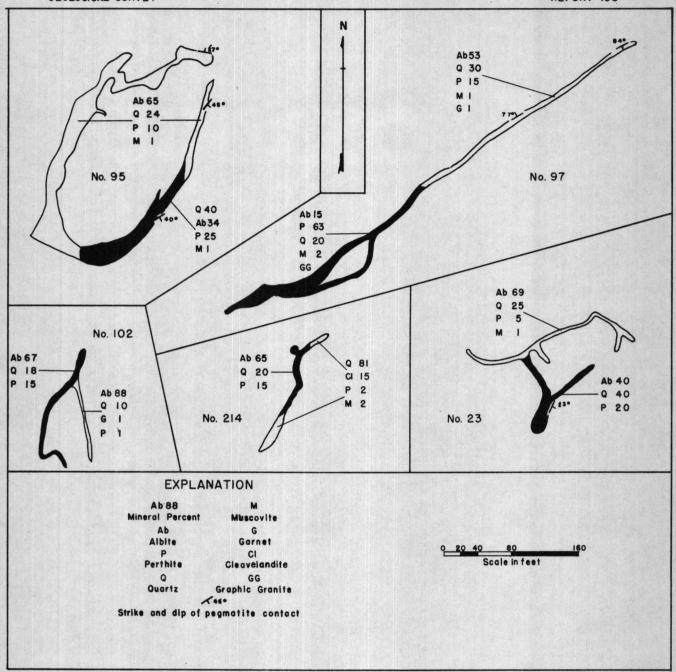


FIGURE 12. PEGMATITES SHOWING VARIATION IN COMPOSITION ALONG STRIKE
QUARTZ CREEK PEGMATITE DISTRICT

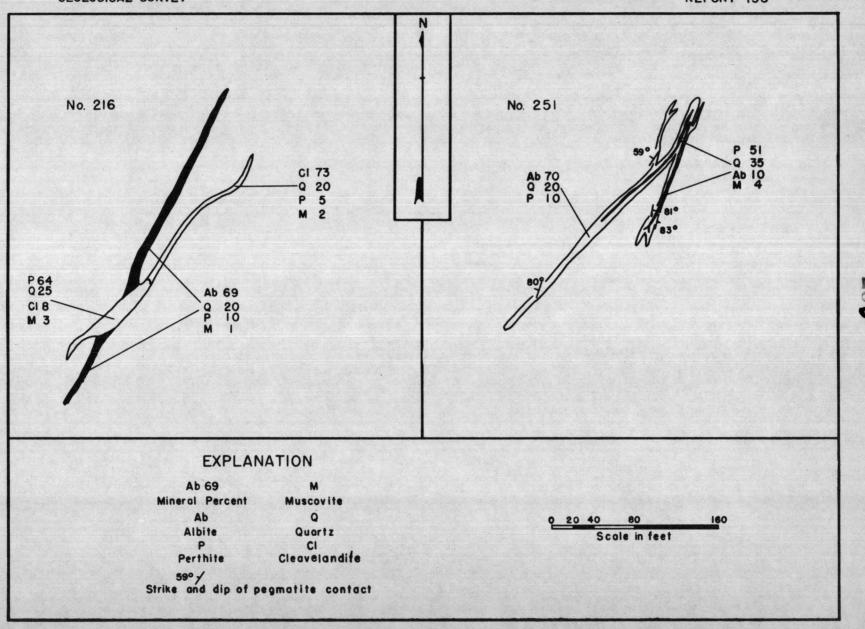


FIGURE 13. MULTIPLE PEGMATITES, QUARTZ CREEK PEGMATITE DISTRICT

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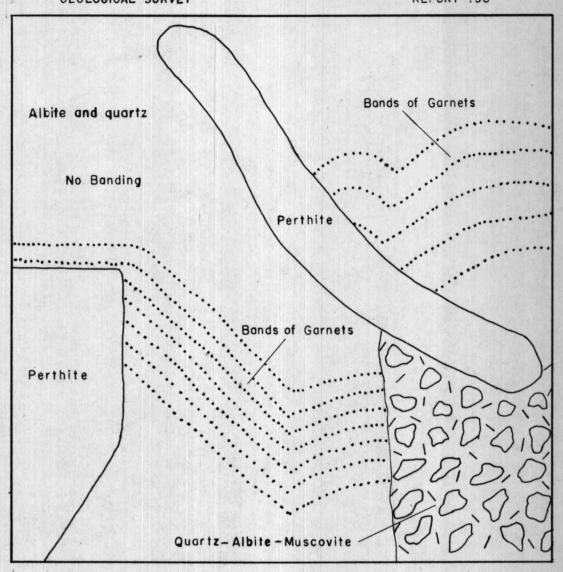
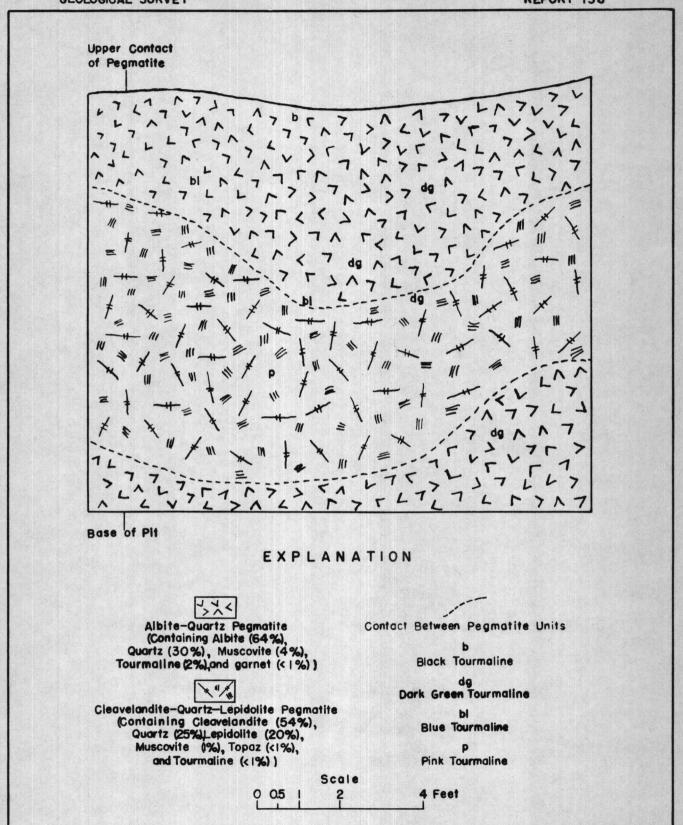


FIGURE 14. RELATION OF BANDING TO CRYSTALS OR NON-BANDED MINERAL AGGREGATES, PEGMATITE NO. 70



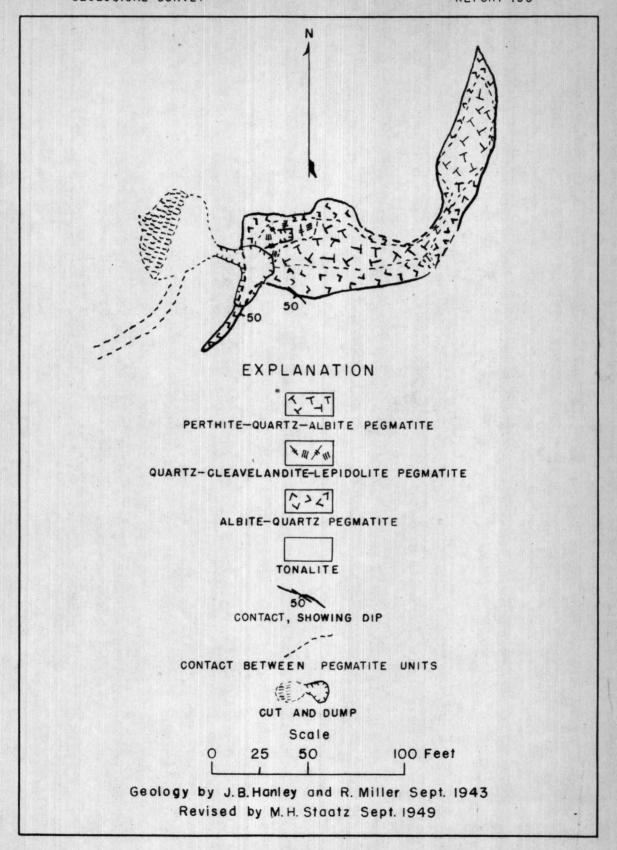


FIGURE 23. GEOLOGIC MAP, BROWN DERBY NO. 5 (NO. 535) PEGMATITE

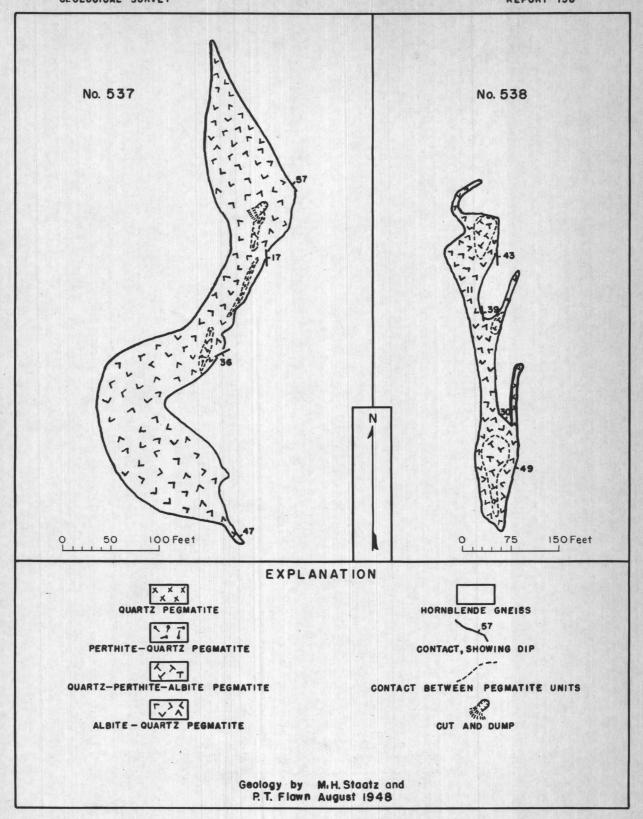


FIGURE 24. GEOLOGIC MAP OF PEGMATITES NO. 537 AND NO. 538

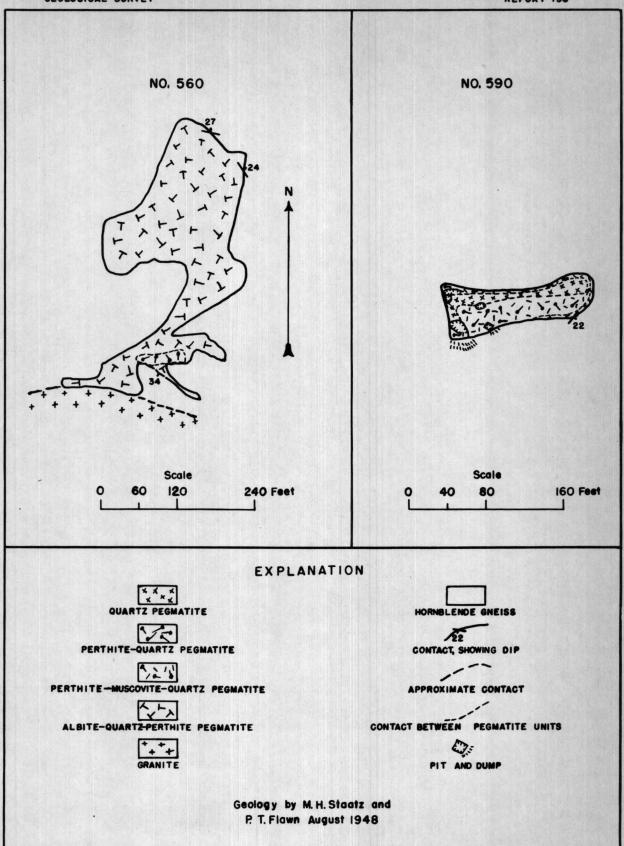


FIGURE 25. GEOLOGIC MAPS OF PEGMATITE 560 AND BERYL AND RARE MINERALS LODE (NO. 590)

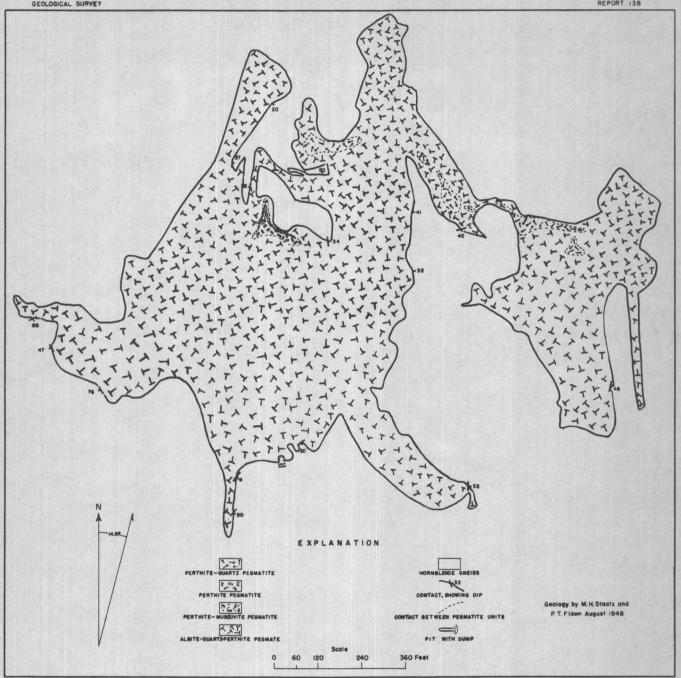


FIGURE 26. GEOLOGIC MAP, BUCKHORN (NO. 659) PEGMATITE

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TEPCO, Washington	4
(Including mester)	45

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METHOD OF CALCULATING PEGMATITE MINERAL RESOURCES

The resources of beryl, potassium feldspar, scrap mica, lepidolite, spodumene, amblygonite, columbite-tantalite, microlite, monazite, and topaz have been calculated on the basis of arbitrary cut-off points. These points were necessary to separate those pegmatites containing small quantities of industrial minerals from those deposits that have possibilities of becoming commercial deposits in the future.

Most of the beryl-bearing pegmatites contain only one or two small crystals of beryl. Beryl-bearing pegmatites that contain less than two square inches of exposed beryl are listed in table 21. This table gives the number of beryl crystals observed and their size range for each pegmatite. Pegmatites with an exposure of more than two square inches of beryl are listed in table 22, with the potential beryl resources of each unit.

The perthite and muscovite, which would help pay for mining the beryl, are calculated in each unit that contains over two square inches of beryl.

Potassium feldspar is found in a large proportion of the pegmatites.

This mineral has a low unit price and, therefore, must be mined in large quantities by inexpensive methods in order to make a profit. To add to the cost of mining is the relatively high transportation costs. A narrow gauge railroad passes through Parlin but the cost of reloading at Salida to a standard gauge car makes it cheaper to truck the 58 miles directly to Salida. The usual method of separating the feldspar from the remainder of the pegmatite is by hand cobbing. In recent years, however, several mills have been erected. The closest mill to the Quartz Creek district is the mill of the Consolidated Feldspar Company at Parkdale, Colo. A

separation is made between hand cobbing and milling feldspar with specifications for each. To be able to hand cob feldspar, the pegmatite must be sufficiently large to give sample tonnages, have crystals large enough to be easily hand cobbed, and be of sufficiently high tenor so as to profitably handle the rock. Limits on these three items are set low enough to anticipate a considerable improvement in mining, cheaper transportation, and a rise in the unit value. The resources of both hand cobbing and milling potash feldspar are given in table 22. The requirements for hand cobbing potash feldspar pegmatites are as follows: a maximum width of the pegmatite unit in excess of 40 feet and a minimum length of from 200 to 300 feet depending on the width; an average grain size of the potassium feldspar in excess of 12 inches; and the tenor of the rock to exceed 25 percent potassium feldspar. All pegmatites falling under this category also have beryl reserves in the Quartz Creek district.

In milling practice grain size is of little importance, and the feld-spar resources of the graphic granite pegmatites are all calculated under this classification. By far the greater feldspar resources in the Quartz Creek district fall under this heading as graphic granite pegmatites are the more common throughout the area. Potash feldspar-bearing pegmatites must meet the following requirements to be considered as milling pegmatites: a maximum width of 40 feet and a length of at least 300 feet, and have a tenor in excess of 15 percent potassium feldspar.

No sheet mica is found in this district and there are only three pegmatite units which have scrap mica in excess of 15 percent. These all contain beryl. Scrap mica reserves are figured, therefore, as byproducts of

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beryl-bearing pegmatites.

The Quartz Creek district is noted for its rarer minerals. Many of these occur in small pods or as one or two crystals. Because of the interest these pegmatites have caused in the past, all pegmatites which contain lepidolite, spodumene, amblygonite, microlite, columbitetantalite, monazite, or topaz are included in table 22.

RESOURCES

A section on reserves in Part I gives the district wide totals of the various economic minerals. Table 22 in Part II gives the reserves broken down to the various properties. Though the reserves of some properties can be published in Part I, the failure to get permission to publish on all properties does not permit a complete listing of properties that is given here. Some pegmatites have only one mineral of economic interest, such as feldspar, but many have several minerals. Reserves are calculated for various units on 134 pegmatites. Only 53 of the total pegmatites on which reserves were calculated were of one unit, 35 of these were feldspar-rich.

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Table 21.—Beryl-bearing pegmatites with less than two square inches of exposed beryl

Pegmatite No.	No. of crystals	Size range in inches	Pegmatite No.	No, of crystals	Size range in inches
21 22 24 55 68 78 80 81 84 86 88 102 105 106 112 114 116 117 121 125 126 128 129 130 131 132 133 134 135 137 139 147 148 150 162 163 172 173 174 184 186	4 1111222124212423223511213222214113128135121	1/8 by 1/4 to 3 by 3 1/4 4 1/2 1-1/2 1/4 to 2 1/8 1/8 1/8 1/8 1/8 1/8 1/8 1/8 1/8 1/8	188 190 196 197 198 201 223 230 231 233 239 244 245 281 283 299 301 302 303 311 316 325 338 342 348 356 363 363 363 363 363 363 363 363 363	2225111563123114355311882231211312222131410161	1/8 1/8 1/16 to 3/4 1/4 1/4 1/2 1/4 1/8 to 1/2 1/8 to 1/2 1/8 to 1/4 1/2 1/8 to 1/4 1/2 1/8 to 1/4 1/8 to 1/4 1/16 to 1/4 1/2 1/2 1/6 to 1/4 1/4 1/2 1/2 1/6 1/4 1/4 1/4 1/6 1/4 1/8 1/4 1/4 1/4 1/4 1/8 1/8 1/4 1/8 1/4 1/8 1/4 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/4 1/8 1/8 1/8 1/8 1/8 1/8 1/8 1/8 1/8 1/8

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Table 21.—Beryl-bearing pegmatites with less than two square inches of exposed beryl—Continued

Pegmatite No.	No. of crystals	Size range in inches	Pegma ite	No. of crystals	Size range in inches
377 378 379 382 384 390 413 416 418 420 425 431 433 435 437 438 439 440 457 463 477 463 477 483 541 551 589 598 602 643 728 905 905 908 909 901 1002	1914112412322t411122111214144112133912234113	1/4 1/4 1/4 1/16 3/4 1/2 1/4 1/4 1/4 1/4 1/4 1/4 1/4 1/4 1/4 1/4	1003 1004 1005 1006 1009 1010 1011 1012 1014 1015 1016 1027 1028 1031 1032 1033 1034 1036 1039 1040 1041 1042 1043 1044	1519261225121112121147133321151124536321653	1/4 1/16 to 3/1 1/4 1/32 to 1/1 1/4 to 1/2 1/16 to 1/1 1/4 3/4 1/4 to 1/2 1/8 1/4 to 3/8 3/16 1/16 to 1/ 3/16 1/16 to 1/ 1/32 to 1/ 1/16 to 3/ 1/8 to 3/1 3/8 to 1/2 1/16 to 3/ 1/8 to 1/4 1/8 1/8 to 5/8 3/16 to 5/ 1/16 to 1/ 1/8 1/8 to 1/4 1/8 to 5/8 3/16 to 5/ 1/16 to 1/ 1/8 to 1/4 1/8 to 3/8 1/16 to 1/4 1/16 to 1/ 1/16 to 3/

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Table 21.—Beryl-bearing pegmatites with less than two square inches of exposed beryl—Continued

Pegmatite	No. of crystals	Size range in inches	Pegmatite No.	No. of crystals	Size range in inches
1086	1	3/16	1174	8	1/4 to 1
1088	3	3/16	1179	8	1/4 to 1/2
1091	1	1/4	1189	1	3/4.
1092	1	1/4	1192	22	1/8 to 1/2
1093	2	1/4	1194	2	1/32 to 1/16
1094	4	1/8 to 1/2	1201	1	1/2
1095	2	1/8	1202	4	5/16 to 1/2
1100	2	1/4	1238	3	3/4 to 1/2
- 1102	1	3/16	1322	5	1/2 to 3/4
1105	1	1/16	1560	1	1-1/2
1172	1	1	1573	1	1

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1105	1	1/16	1560	í	1-1/2
1172	1	i	1573	1	1

205 Core. 213 Wall sone. Opportunity No. 1 215 Central unit Opportunity No. 1 216 South- Opportunity No. 1 216 Wortheast Opportunity No. 1 216 Central and Opportunity No. 1 216 Central and Opportunity No. 1 217 South half wall sone. Opportunity Nos. 22 and 34 250 Core pods. Opportunity Nos. 22 and 34	ructure	Shape	Average length (feet)	Average thick- ness (feet)	Depth	Tons	Per-	Сопро-	Tons						3.	1			-			T		LI		T
23 Morth brand core. 64 Vest branch core. 136 One unit. 205 Core. 213 Wall zone. Opportunity No. 1 215 Central unit. Opportunity south- Opportunity south No. 1 216 Mortheast Opportunity branch. No. 1 216 Central and Opportunity wall zone. Opportunity No. 22 250 Vall zone.		Lantiania						sition (Percent BeO)		Per- cent cobb- able	Percent	Compo- sition	Tons	Per- cent cobb- able	cent	punch	Size of sheet and punch	Quality of sheet and punch	sheet	Total tons of scrap	Total tons of mica	Per- cent cobb- able	Mineral	Per- cent	Compo- sition	To
core. 136 One unit. 205 Core. 213 Wall zone. Opportunity No. 1 215 Central unit. Opportunity No. 1 216 Wortheast Opportunity No. 1 216 Central and Opportunity No. 1 217 Central and Opportunity No. 24 250 Wall zone. 250 Cere pods. Opportunity Nos. 22 250 Core pods.	th branc	Tention19	. 155	9																			Columbite- tantalite.	13		
core. 136 One unit. 205 Core. 213 Wall zone. Opportunity No. 1 215 Central unit Opportunity No. 1 216 Wortheast Opportunity No. 1 216 Central and Opportunity No. 1 216 South Opportunity No. 1 216 Wortheast Opportunity No. 1 216 Central and Opportunity No. 1 245 South half wall zone. Opportunity Nos. 22 and 34 250 Core pods. Opportunity Nos. 22 and 34	re.	h do.	165	2 to 4	25	1060	0.03	13.1	0.32	95	55		590	100	0	0			0	20	20	50	Monazite. Columbite- tantalite.	1 ₃ 1 ₅		
213 Vall sone. Opportunity No. 1 215 Central unit Opportunity No. 1 215 South- Opportunity No. 1 216 Wortheast Opportunity No. 1 216 Central and Opportunity No. 1 216 Central and Opportunity No. 1 217 South half wall sone. Opportunity No. 22 and 34 250 Core pods. Opportunity Nos. 22 and 34		do.	50	1 to 3						12.1													fonazite. Columbite- tantalite.	1 ₁		
Opportunity No. 1 215 Opportunity No. 1 215 Opportunity No. 1 216 Opportunity No. 1 216 Opportunity No. 1 216 Opportunity No. 1 245 South half wall sone. 250 Opportunity Nos. 22 and 3A 250 Core pods. Opportunity Nos. 2A and 3A		Irregular.	195	20	25	7160	0.002	13.2	0.14	30	8		570	0									Columbite-	12		
215 Central unit Opportunity No. 1 215 Opportunity No. 1 216 Opportunity No. 1 216 Opportunity No. 1 216 Central and Opportunity No. 1 245 South half wall sone. 250 Opportunity Nos. 24 and 34 250 Core pods. Opportunity Nos. 24 and 34	l sone.	do.	165	10	25	3580	0.08	13.0	2.8	95	12		430	50									tantalite.			
Opportunity central unit. 216 Unit branch. No. 1 216 Central and opportunity branch. 245 South half wall sone. 250 Opportunity Nos. 24 and 34 250 Core pods. Opportunity Nos. 24 and 34	ral unit	Lenticula branching.	r-250	14	25	7500	0.1	13.3	7.0	90	8.0		600	50	0	0			0	150	150	0	Columbite- tantalite.	0.5 9.1 21 16 14		
Opportunity branch. 216 Opportunity southwest- No. 1 245 South half wall sone. 250 Opportunity Nos. 24 and 34 250 Core pods. Opportunity Nos. 24 and 3A	ntral	Irregular.	90	27	25	5170					60		3100	75									Monazite.	-4		
Opportunity southwest- No. 1 south helf wall sone. 250 Wall sone. Opportunity Nos. 2A and 3A Core pods. Opportunity Nos. 2A and 3A	theast anch.	Lenticular	. 167	7	25	2580	0.01	13.5	0.26	75	5		130	50	0	0			0	50	50	5	Microlite.	12		Child agenting
vall sone. 250 Wall sone. Nos. 24 and 34 250 Core pods. Opportunity Nos. 24 and 34	uthwest-		106	12	25	2770	0.005	13.3	0.14	95	64		1770	100	0	0			0	80	80	0				
Opportunity Nos. 2A and 3A 250 Core pode. Opportunity Nos. 2A and 3A	th half	Irregular	1250	450																			Columbite- tantalite.	13		
Opportunity Nos. 2A and 3A	l sone.	Lenticular branching		35			11	13.0										4					Monazite.	1,		
	ore pode	Lenticula	. 248	11	20	4870	0.02	13.1	1.0	100	30		1460	80												
251 Core.	ore.	do.	120	0.6	20	125	0.03	13.0	0.04	50	51		64	50	0	0			0	5	5	10	Columbite- tantalite. Monazite.	1 ₂		
266 do.	do.	do.	26	4							500												Columbite- tantalite.	1,		
270 do.		do.	90	82	20	310	0.02	13.2	0.06	50	30		90	60												

1/ Humber of crystals observed.

			ize and	shape of	depos	lt		Beryl			And a	Feld	par					Mica					Ot	her mi	nerals	
Number and name of peg- matite (Pl.II)	Internal structure	Shape	Average length (feet)	Average thick- ness (feet)	Depth (feet)	Tone	Per- cent	Compo- sition (Percent BeO)	Tons	Per- cent cobb- able	Per- cent	Compo- sition	Tons	Per- cent cobb- able	Per- cent crude sheet and punch	punch	Size of sheet and punch	Quality of sheet and punch	sheet	Total tons of scrap	Total tons of mica	Per- cent cobb- able	Mineral	Per- cent	Compo- sition	1
289	Core.	Lenticula	. 90	2	10	150	0.02	13.2	0.03	80	5		7.5	60									Columbite- tantalite.	11		
290	do.	do.	1-1/2	1-1/2																			Monasite.	16		1
290	do.	do.	20	1						2100													Columbite- tantalite.	11		
306 Opportunity No. 4	West layer.	do.	24	3	10	60									0	0			0	0.6	0.6	0	Lepidolite. Microlite.	15		
358	One unit.	do.	95	12	25	2480	.002	13.1	0.05	0	20		500	5												
359	do.	do.	257	14	25	7820	0.0008	13.1	0.06	50	20		1560	5									Columbite- tantalite.	11		
385	do.	do.	79	6								T.											do.	12		1
411	Core.	do.	60	30	25	3920	0.9		0.39	50	35		1370	80	0	0			0	200	200	0				1
417	do.	Irregular	360	2	25	1920	0.2	13.3	3.8	60	45		860	80	0	0			0	96	96	0				1
H22 Ben-Kauf No. 1	do.	Lenticula	. 6	5	25	65					81		352	80	0	0			0	9.8	0	0	Lepidolite. Microlite.	2 0.01	-	1
424 Bazooka	Wall zone.	do.	228	10	25	4950					85		34060	70									Lepidolite.	8	1	1
424 Basooka	Core.	Oval.	20	20	25	800					20		3160	90				*					do. Microlite. Spodumene. Amblygonite	1 ₁ 12 6		
432	Hanging wall layer.	Lenticula	. 660	6	25	8300	0.05	13.2	4.2	30	30		2490	60	0	0			0	330	330	0	Z=01/gonito			
H. ADAMS S.	Foot wall layer.	do.		-			12-4	13.4																		
434	One unit.	do.	32	7	20	390	0.1	13.4	0.4	100	8		30	20	0	0			0	10	10	0				1
H52 Brown Derby Dike No. 1	West branch core.	do.	20	0.5	8																		Columbite- tantalite. Monazite.	1.4		2 2
Brown Derby Dike No. 1 Perthite- quartz- albite pegmatite.	Hanging wall layer	do.	785	8	140	76,450	0.1		76	95	140		30,580	100	0	0			0	7,600	7,600	0				
452 Brown Derby Dike No. 1	Curved lepidolite layer.	do.	190	2	140	4,630					1414		32040	90									Lepidolite. Topaz.	15		
452 Brown Derby Dike No. 1	Lepidolite microlite pod.	do.	50	6	10	260	11																Lepidolite. Microlite. Topaz.	40 .03		

Number of crystals observed. Founds. Soda spar.

		S	ize and	shape of	deposi	lt		Beryl		-	100	Feld	spar		1000			Mica					Ot	her mi	nerals	
Aumber and name of peg- natite (Pl. II)	Internal structure	Shape	Average length (feet)		Depth (feet)	Tons	Percent	Composition 5/ (Percent BeO)	Tons	Per- cent cobb- able	Per- cent	Compo- sition	Tons	Per- cent cobb- able	cent crude sheet and punch	sheet and punch	Size of sheet and punch	Quality of sheet and punch	Tons sheet and punch	Total tons of scrap	Total tons of mica	Per- cent cobb- able	Mineral	Percent	Compo- sition	Ton
Brown Derby Dike No. 1			319	2	140	7.770	14				25		31940	90									Lepidolite. Topaz.	10		777
Brown/Derby Dike No. 2	Core.	do.	180	1.5	240	5,640					53		32990	70									Lepidolite. Topaz. Kicrolite.	5 1111		28
H5H Brown Derby Dike No. 3	do.	do.	108	1	150	1410	0.3		4.2	95	1 47		3660	100 80	0	0			0	56	56	0	Lepidolite. Columbite- tantalite. Microlite. Topas.	5 11 14 4		71
H55 Brown Derby Dike No. 4	Southern part, core.	do.	342	8	25	5750	0.076	13.2	4.4	100	80		4600	100						*						
456 Brown Derby Dike No. 5	Core.	do.	50	1																			Microlite.	11		
457 Brown Derby Dike No. 6	Wall zone.	do.	175	12	25	2600																	Lepidolite.	3		78
457 Brown Derby Dike No. 6	Core.	do.	120	4	25	1040																	do.	1		10
458 Brown Derby Dike No. 7	Hanging wall layer	do.	95	5	25	1030	.007	13.2	0.07	80	70		720	70								7				
H59 Brown Derby Dike No. 8	Wall zone.	Lenticular branching.		65	25	86,200	.001	13.4	0.86	80	35		30,200	85	0	0			0	2590	2590	10				
Brown Derby Dike No. 8	Core pod.	Lenticular	. 3	1	10	2.6					67		31.7	80	0	0			0	0.08	0.08	10	Lepidolite.	2 10		0.0
H61 Brown Derby Dike No.10	One unit.	Lenticular branching.	- 560	50	25	54,900	.0006	13.3	0.33	50	35		19,200	50	0	0			0	550	550	0				
462	Core.	Lenticular		65	50	236,000					65		¹¹ 153,000	10									Microlite. Lepidolite.	15 <1		
498	do.	do.	300	105	50	149,000					35		452,000	20												
503	do.	Lenticular branching.	This is	40	50	64,000					65		411,600	10	0	0			0	1920	1920	0				
505	Main unit.	Lenticular	. 480	. 60	50	90,900					45		440,900	10	1				1		-					

^{1/} Number of crystals observed. 2/ Pounds.

^{3/} Soda feldspar. 1/ Mainly graphic granite.

		S	ize and	shape of	deposi	lt		Beryl				Feld	spar					Mica					Ot	her mi	nerals	
umber and ame of peg- atite Pl. II)	Internal structure	Shape	Average length (feet)	Average thick- ness (feet)	Depth (feet)	Tons	Percent	Composition 5/ (Percent BeO)	Tons	Per- cent cobb- able	Percent	Compo- sition	Tons	Per- cent cobb- able	Per- cent crude sheet and punch	punch	and	Quality of sheet and punch	Tons sheet and punch	Total tons of scrap	Total tons of mica	Per- cent cobb- able	Mineral	Percent	Compo- sition	Tor
509	Main unit.	Irregular	. 600	60	50	158,000					30		47,000	80												
512	One unit.	do.	780	85	50	246,000					39		496,000	5	0	0			0	2500	2500	0				
516	do.	Lenticular	. 480	48	50	77,000	4				35		427,000	10	0	0			0	770	770	0				
517	Main unit.	do.	300	45	50	56,000					25		14,000	15	-11											
534	Core.	do.					1,																			
535 Brown Derby No. 5	Wall zone.	Irregular	. 207		25	6200	.01	13.3	0.62	30	10		620	10									Lepidolite	Trace.		
535 Brown Derby	Intermediat zone.	e Lenticul	ar.33	13	20	750	0.1	13.05	0.75	80	14 35		3,30	90 70	0	0			0	7.5	7.5	20	Lepidolite. Topaz. Microlite. Columbite- tantalite.	5 121 113 11		37
535 Brown Derby No. 5	Core.	Irregular.	175	13	50	4,000	.003		0.12	100	39		1560	95	0	0			0	40	40	10	Lepidolite. Columbite- tantalite.	Trace.		
537 Brown Derby No. 4	do.	Lenticular	. 52	10	25	1125	0.31	13.4	3.4	65	66		745	80	0	0			0	45	45	5				
538	do.	Irregular.	128	35	25	4360	0.95	13.4	42	75	32		1400	80	0	0			0	131	131	0				
548	Hanging wall layer.	Lenticula	r.125	7	25	1900	.016	13.4	0.30	95	50		950	30	0	0			0	38	38	0				
550	One unit.	do.	312	20																			Monazite.	11		
557	Wall zone.	do.	1110	70	50	313,000					20		62,600	0	0	0			0	3130	3130	0				
558	Main unit.	Lenticula branching.	r-1044	75	50	707,000					20		141,400	0	0	0			0	7070	7070	0				
560	Core pod.	Lenticular	- 74	20	15	2120	0.45	13.4	9.5	100	55		1170	100					196							
564	Main unit.	Irregular	. 600		50	374,000					20		74,800	0	0	0			0	3740	3740	0				
587	One unit.	do.	720	135	50	352,000					20		70,400	0	0	0			0	3520	3520	0				
590 Beryl and Rare Minera Lode	Intermediat zone.	e Lenticul	ar.155	41	3	1450	0.1		1.5	85	50		725	100	0	0			0	465	465	80	Columbite- tantalite.	.002		2
596	One unit.	do.	520	50	50	113,000					40		45,200	50	0	0			0	1130	1130	0				
604 White Spar No. 1	Intermediate	do.	90	6.5	25	1300	0.01		0.13	65	15 50		194 3650	85 50									Topas. Lepidolite. Columbite- tantalite. Microlite.	.003 5 .0003		2 24
604 White Spar No. 1	Pod.	do.	7	3.3	1	2																	Lepidolite. Microlite.	90 0.1		24

1/ Number of crystals observed. 2/ Pounds. 3/ Soda spar.

4/ Mainly graphic granite.

5/ Composition of beryl determined by index of refraction according to a chart by W. T. Schaller (unpublished).

		S	ize and	shape of	deposi	t		Beryl		_	100	Felds	per					Mica					Ot	her mi	nerals	1986
umber and ame of peg- atite (Pl. II)	Internal structure	Shape	Average length (feet)	Average thick- ness (feet)	Depth (feet)	Tons	Percent	Composition (Percent BeO)	Tons	Per- cent cobb- able	Per- cent	Compo- sition	Tons	Per- cent cobb- able	Per- cent crude sheet and punch	and	Size of sheet and punch	Quality of sheet and punch	Tons sheet and punch	Total tons of scrap	Total tons of mica	Per- cent cobb- able	Mineral	Percent	Compo- sition	To
605	One unit.	Lenticular	. 100	3	25	720					25		180	25	0	0			0	7.2	7.2	0	Lepidolite.	2		14
607	Fracture filling.		Small.								23			30									do.	2		
623	Pod.	Lenticular	. 14	4	15	73	0.6		0.44	35	40		29	30												
636 White Spar No. 2	Wall zone.	Irregular.	130	14	25	2360																	Lepidolite.	<1		
636 White Spar No. 2	Core.	Lenticular	. 235	24	25	9820	12				45		3141400 980	0									do. Microlite.	130		9
637	do.	Irregular.	147	45								*	,										Lepidolite. Golumbite- tantalite.	Trace	•	
637	Fracture filling.	Lenticular	- 55	6			11																Lepidolite.	Trace		
659 Buckhorn	Intermediat	e do.	145		1.5	600	14			20	50		300	100	0	0			0	240	540	95	Columbite- tantalite. Monazite.	14		
659 Buckhorn	Core, Buckhorn claim.	do.	145	-	5.5	2200	0.7		15.4	98	91		2000	100	0	0			0	22	22	0				
699 Buckhorn	Core pods,	Irregular.	_		15	23,960	.002		0.50	70	53		12,700	100	0	0			0	240	240	0				
661	Core.	Lenticular	. 10	1.5	10	13	8.5		1.1	85	42		5.5	95	0	0			0	1.3	1.3	0				
671	Main unit.	Irregular.	690		50	738,800					19		4140,400	0	0	0			0	7400	7400	0				
697	One unit.	Lenticular	. 350	50	25	36,500					54		419,700	0	0	0			0	365	365	0				
783	Wall zone.	Irregular.	2300	70	50	1,380,000					50		4690,000	0						- *				75		
783	Core pod.	Lenticular	. 8	2	10	14	1.2		0.17	100	65		9.1	100	0	0			0	0.6	0.6	0				
790	One unit.	do.	640	55	25	55,200					65		435.900	0												
826	Wall sone.	do.	24	14																			Monazite.	Trace		
827	One unit.	do.	30	10						100													do.	Trace		
828	do.	do.	24	4						1													do.	Frace.		
837	do.	Irregular.	360		50	254,000		Sec. 17.			65		4165,000	0										lane!		
g47 Black Wonder	Wall zone	do.	13,600	-	350	543,400,0	0				35		¹ 225,200,	000 0												
847 Black Wonder	Intermedia sone.	te Lenticul	ar. 30	10	25	980	0.2	13.4	2.0	60	5		49	50	0	0			0	49	149	40				

1/ Number of crystals observed.
3/ Soda spar.

4/ Mainly graphic granite.

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			ize and	shape of	depos	dt		Beryl	_	1916		Feld	par					Mica					Ot	her mir	nerals	
fumber and name of peg- natite (Pl.II)	Internal structure	Shape	Average length (feet)	Average thick- ness (feet)	Depth (feet)	Tons	Percent	Composition (Percent Be0)	Tons	Per- cent cobb- able	Per- cent	Compo- sition	Tons	Per- cent cobb- able	Per- cent crude sheet and punch	Percent sheet and punch in crude	Size , of sheet and punch	Quality of sheet and punch	Tons sheet and punch	of	Total tons of mica	Per- cent cobb- able	Mineral	Percent	Compo- sition	Fons
847 Black Wonder	Intermediat zone.	e Lenticul	ar. 15	14	10	52									0	0			0	16	16	85	Monasite. Columbite- tantalite.	.007		293
852	One unit.	Irregular	. 350	160	50	213,000					45		495,850	10												
855	do.	do.	800		50	472,000					65		4306,800	5												
860	do.	do.	364	90	50	115,400		A STATE			50		457.700	10								19.00				
963	Hanging wall layer	Lenticular	. 165	4																			Monasite.	11		
985	do.	do.	20	2							- 7												do.	13		
997	Foot wall layer.	do.	165	6																			do.	11		
1002	Core.	do.	20	1																			Columbite- tantalite.	11		
1035	One unit.	do.	55	2						15.0				6.56									do.	11		
1036	do.	do.	100	12																			do.	1,		
1049	do.	do.	40	2																			do.	1,		
1064	do.	do.	470	10																			do.	1,		
1068	Core.	do.		2											- X								do.	11		
1140	One unit.	Irregular	440	100	50	153,000					25		38,250	5	0	0			0	3060	3060	0				
1142	do.	enticular	520	120	50	266,000	7/0				35		93,100	10												
1150	Core.	do.	3	1.5	3	1.2	2.8	13.4	0.03	100	30		0.4	80	0	0			0	0.06	0.06	0				
1182	do.	do.	50	1	25	109	0.46		0.50	100	5		5.5	90												
1193	do.	Oval.	6	4	5	10	0.12	13.3	0.01	80	20		2	90												
1234	Wall zone.	enticular	176	14				X7 (1)															Columbite- tantalite.	10		
1240	One unit. I	enticular branching	50	11																*			do.	11		
1248	Wall sone.	Irregular	2335		50	2,450,000			1		20		4490,000	0												
1256	One unit.	Lenticula	. 500	35	25	37.400					30		411,200	0												
1269	do.	Irregular	410	140	50	240,000					25		60,000	0	SE AS						1					
1316	do.	do.	360	65	50	108,500					35		439,600	5												
1322	Wall zone.	Lenticular branching.	- 2780		50	4,260,000					30		41,278,00	0												
1323	Core.	Oval.	15	15	10	196	0.2		0.39	100																
1332	One unit.	Irregular	568	65	50	252,000	1823		1000	10.3	25		63,000	30	8501	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1			1000		1	1400	The second	-		1

280

1/ Number of crystals observed. 2/ Pounds.

4/ Mainly graphic granite.

	Waste of		ize and	shape of	depos	1t	1	Beryl		_		Feld	spar		1000	Diam.		Mica	(30)	To the letter			0	ther mi	nerals	
Number and name of peg- matite (Pl. II)	Internal	Shape	Average length (feet)	Average thick- ness (feet)	Depth (feet)	Tons	Percent	Composition (Percent BeO)	Tons	Per- cent cobb- able	Per- cent	Compo- sition	Tons	Per- cent cobb- able		and	Size of sheet and punch	Quality of sheet and punch	Tons sheet and punch	of	Total tons of mica	Per- cent cobb- able	Mineral	Percent	Compo- sition	Ton
1341	Hanging wall layer	Lenticular	. 580	20	50	52,700					45		423.700	5												
1341	Foot wall layer.	do.	880	45	50	168,700		No.			20		433,700	0												
1347	One unit.	Irregular.	314	38	25	39,000					30		411,700	10												
1355	Wall sone.	do.	3,450		90	7,826,000					35		42,739.00	0 15												
1357	One unit.	do.	535	60	50	185,000					25		446,200	0												
1402 Trio No. 1	Wall zone.	do.	644		50	290,000					25		472,500	0												
1402 Trio No. 1	Intermediat	e Lenticul	ar.96	5	20	860	0.91	13.2	7.5	95	30		258	100												
1426	Wall mone.	Oval.	330	115	50	135,500					70		94,900	5	6.63											
1428	do.	Lenticular branching			50	262,000					65		170,000	0												
1429	do.	Irregular.	385	76	50	109,000			Man.		45		1419,000	0					B WH							
1470	One unit.	do.	1,140	-	50	783,000	100		Deg Cit		30		4235,000	5	111/2/112							Total Co				
1509	do.	do.	1,000		50	1,226,000					48		588,000	30	0	0			0	24,520	24,520	5				
1517	do.	Lenticular	. 34	12					lates.														Monazite.	11		
1518	do.	Lenticular	•	20																			do.	11		
1523	Wall sone.	Irregular.	600		50	379,000					30		4113,700	P.C. P.S. C.C.												
1534 The Trio	do.	do.	2,800		100	2,287,000					30		14686,000													
1540	do.	do.	500	190	50	289,100					33		495.400	5	0	0			0	5,780	5,780					
1541	do.	do.	1,020	260	50	1,217,000					30		365,000	25	0	0			0	60,850	60,850	0				
1544	do.	do.	368	170	50	242,000					50		4121,000	5	0	0			0	2,420	2,420	0				
1557	Core.	Lenticular	. 24	6		100								16									Columbite- tantalite.	11		1
1558	do.	do.	种	1.5																			do.	1,		000
1566	One unit.	Irregular.	384	75	50	131,300					30		39.400	0	0	0			0	1,310	1,310	0				100
1574 Bucky	Graphic granite, Intermediat zone.	do.	1,350	650	200	16,870,00	d				80		413,500,0	00 0												S den
1574 Bucky	Muscovite- feldspar quartz-ber intermedia sone.		100	90	3	1.770	8.9	12.9	158	100	26		460	100	0	0			0	708	708	100	Lepidolite. Topaz. Columbite- tantalite. Monazite.	Trace.		1.95

Mainly graphic granite.

		S	ize and	shape of	depos	11	1	Beryl				Feld	spar		-			Mica					Ot	her mi	nerals	
nber and ne of peg- tite 1.)	Internal structure	Shape	Average length (feet)	Average thick- ness (feet)	Depth (feet)	Tons	Percent	Composition (Percent BeO)	Tons	Per- cent cobb- able	cent	Compo- sition	Tons	Per- cent cobb- able	Per- cent crude sheet and punci	cent sheet and punch	Size of sheet and punch	Quality of sheet and punch	sheet	Total tons of scrap	Total tons of mica	Per- cent cobb- able	Kineral	Percent	Compo- sition	Ton
1574 Bucky	Perthite intermediat	Oval.	90	50	17	5200	<1				93		14800	100												
1584	One unit.	Lenticular	. 450	70	50	120,500					55		466,300	0												
585	do.	do.	450	42	50	86,300					75		464,700	0												
589	do.	Irregular.	800	400	50	869,600					59		513,000	0	0	0			0	8700	8700	0				
610	Wall zone.	do.	1110	120	50	646,700					55		4355,700	5							1					
517	One unit.	do.	410	110	50	150,500			100		60		90,300	5						*						
19	do.	do.	651	170	50	426,900					60		4256,000	0												
27	do.	do.	173	89																			Lepidolite.	Trace.		1
530	Wall zone.	Lenticular branching	- 258	7																			Monazite.	11		
57	One unit.	Irregular	. 732	55	50	198,100					55		4109,000	0								200				
67	do.	do.	426	105	50	179,500					55		98,700	0												
100	do.	do.	1820	350	50	2,767,000					55		1,522,000	5												
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1/ Number of crystals observed.
4/ Mainly graphic granite.