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GROUND-WATER RESOURCES OF THE UPPER WINOOSKI RIVER BASIN, VERMONT

U.S. GEOLOGICAL SURVEY

Water-Resources Investigations 77-120

Prepared in cooperation with the STATE OF VERMONT, AGENCY OF ENVIRONMENTAL CONSERVATION, DEPARTMENT OF WATER RESOURCES





GROUND-WATER RESOURCES OF THE UPPER WINOOSKI RIVER BASIN, VERMONT

By

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UNITED STATES DEPARTMENT OF THE INTERIOR

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GEOLOGICAL SURVEY

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Open-File Report

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U.S. Geological Survey 150 Causeway Street, Suite 1001 Boston, MA 02114 GROUND-WATER RESOURCES OF THE UPPER WINOOSKI RIVER BASIN, VERMONT (WRI 77-120)

ERRATA

Plates 1 - 4 Errors in registration of overlays with base maps during final printing are responsible for slight mislocations of the information as shown.

Plate 3 The lense shaped map unit surrounding test site MEX-1 in Marshfield should have the pattern of an area underlain chiefly by coarse-grained saturated stratified deposits.

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CONVERSION FACTORS

The following factors may be used to convert the U. S. customary units published herein to the International System of Units (SI):

Multiply U.S. customary units	Ву	To obtain SI units
	Length	
inches (in)	25.4	millimeters (mm)
	.0254	meters (m)
feet (ft)	.3048	meters (m)
miles (mi)	1.609	kilometers (km)
	Area	
square miles (mi ²)	2.590	square kilometers (km ²)
	Volume	
gallons (gal)	3.785	liters (L)
	3.785×10^{-3}	cubic meters (m^3)
million gallons (10 ⁶ gal)	3785	cubic meters (m^3)
	3.785×10^{-3}	cubic hectometers (hm ³)
	Flow	
gallons per minute (gal/min)	.06309	liters per second (L/s)
	.06309	cubic decimeters per second (dm ³ /s)
	6.309×10^{-5}	cubic meters per second (m^3/s)
million gallons per day (Mgal/day)	43.81	cubic decimeters per second (dm ³ /s)
	.04381	cubic meters per second (m^3/s)

GROUND-WATER RESOURCES OF THE

UPPER WINOOSKI RIVER BASIN, VERMONT

By Arthur L. Hodges, Jr., Richard E. Willey, James W. Ashley, and David Butterfield

ABSTRACT

Ground water in the upper Winooski River basin occurs in bedrock, and in overlying unconsolidated deposits of glacial origin. Bedrock in the area is composed of a series of metamorphic and igneous rocks. Median yield for 126 wells in four different bedrock formations ranges from 5 to 6 gallons per minute (0.32 to 0.38 liters per second), and median depth ranges from 130 to 200 feet (40 to 61 meters). Lineaments, interpreted as fracture or breakage zones in bedrock, were mapped to identify zones where well yields are expected to be higher than average.

Unconsolidated deposits in the upper Winooski River basin include unsorted till, and water-sorted clay, silt, sand, and gravel. Properly constructed wells in saturated deposits of sand or gravel having high permeability can yield large quantities of water. Twenty-six domestic wells in these unconsolidated deposits have a median yield of 18 gallons per minute (1.1 liters per second) and a median depth of 58 feet (17.6 meters).

Chemical analyses of water from six wells in the upper Winooski River basin indicate a median hardness of 120 milligrams per liter (as $CaCO_3$), which is moderately hard. Iron and manganese are common constituents of ground water in the area, and several analyses show concentrations of these elements which exceed recommended National Academy of Sciences and National Academy of Engineering (1973) limits for public drinking water supplies.

INTRODUCTION

Purpose of Investigation

This report describes the ground-water resources and related geologic environment of the upper Winooski River basin and is a part of a continuing program to locate and evaluate ground-water resources in Vermont. The program, begun in 1965, is a cooperative study program carried on by the U.S. Geological Survey and the State of Vermont. Field work for the investigation was started during the summer of 1972 and continued through October 1975.

The purpose of the work was as follows: to determine the thickness and extent of the water-bearing unconsolidated deposits, to evaluate the hydraulic properties of these water-bearing materials and their potential yield, to evaluate the potential yield of bedrock aquifers and to evaluate the variations and concentrations of chemical constituents in the ground water and their effect on its general use.

Location and Extent of Area

The 203 mi² (526 km²) area of investigation includes all or parts of the Towns of Cabot, Calais, Marshfield, Woodbury, and Worcester in the northeast corner of Washington County (fig. 1) and lies adjacent to the Barre-Montpelier area, a major commercial, industrial, and governmental center.

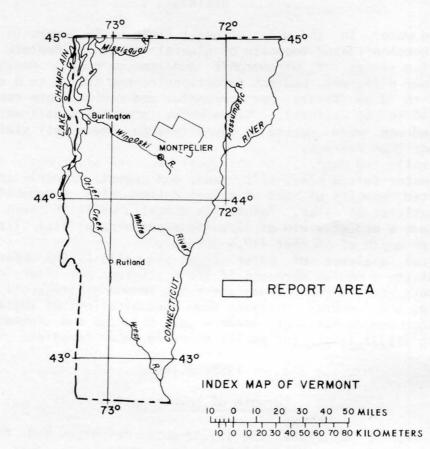


Figure 1.--Location of study area

Previous Investigations

One ground-water study and a number of geologic reports, mostly at the reconnaissance level, have been made covering all or part of the study area. The Department of Water Resources of the State of Vermont in cooperation with the U.S. Geological Survey has published a Ground Water Favorability Map of the Winooski River Basin, Vermont (Hodges, 1967). Geologic studies include bedrock mapping and reports by Cady (1956), Hall (1959), Doll and others (1961), and Konig (1961); and surficial mapping by Doll and others (1970), and Stewart and MacClintock (1969). A more recent report (Stewart, 1971) provides general information on geology for environmental planning.

Methods of Investigation

Field mapping was undertaken to delineate and identify unconsolidated materials, their origin and character, and to map bedrock exposures. Seismic refraction profiling was carried out to determine the thickness of unconsolidated deposits in various areas where water saturated coarse grained materials appeared sufficiently thick to yield large quantities of water to wells. Location maps and seismic refraction profiles for 10 sites tested are shown on plate 4 and in figures 2 and 3.

Based on the results of the seismic investigations, wash borings were made at 23 sites to obtain detailed lithologic descriptions of the underlying unconsolidated deposits. These borings were made during the summers of 1973 and 1974 and involved a total of about 1,200 linear feet (360 m).

Location, construction, and hydrologic data for 186 privately owned water wells were obtained. Most of these wells were drilled into bedrock and provided information on thickness and character of the unconsolidated deposits, and the nature and yield of the bedrock aquifer.

Water-quality analyses of surface water from 23 sampling sites were made by the Vermont Department of Water Resources between 1953 and 1965. Thirteen drilled wells from which water had been chemically analyzed by the Vermont Department of Health were located in the field.

Aerial photographs were examined for geologic information with particular emphasis on the character and location of surficial material, borrow pits or quarries, and lineaments.

The stratigraphic nomenclature used in this report is that used by the Vermont Geological Survey and does not necessarily follow the usage of the U.S. Geological Survey.

Well-Numbering System

Data on wells and borings used in compiling this report are given in tables 6 and 7. Well and boring identification numbers used in this report are composed of a two-letter town code, the letter "W" or "X", and a sequential number. The letter "W" designates a well while "X" represents a test boring. The letter codes for the five towns in the upper Winooski River basin are: Cabot, CA; Calais, CB; Marshfield, ME; Woodbury, X6; and Worcester, X9.

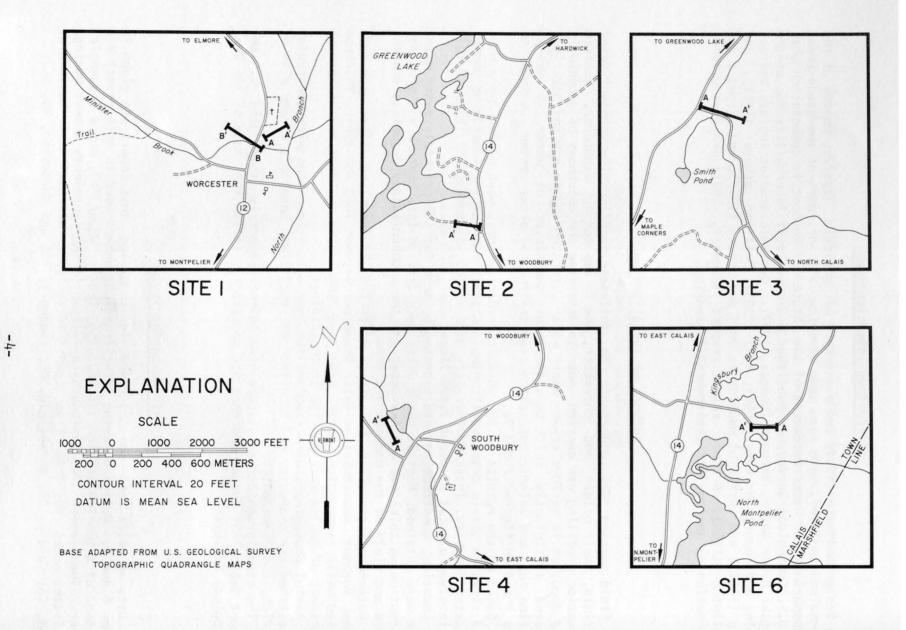
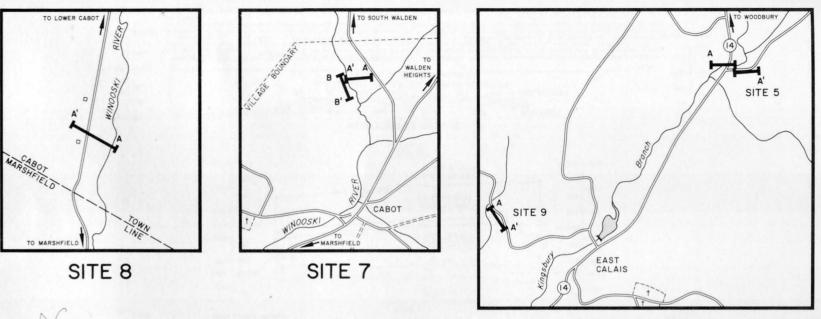
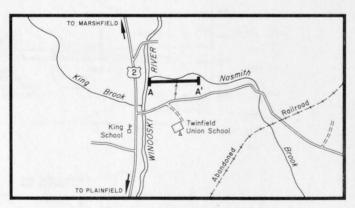


Figure 2 .-- Detailed location maps of seismic surveys



SITES 5 AND 9



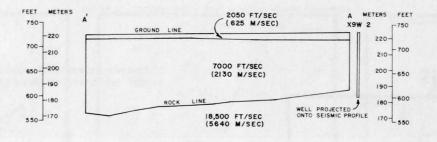
SITE IO

EXPLANATION

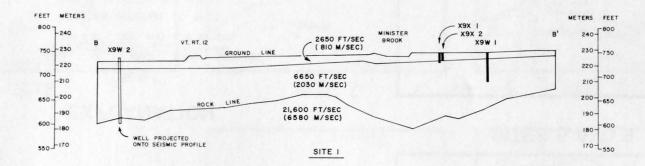
SCALE 1000 0 1000 2000 3000 FEET 200 0 200 400 600 METERS CONTOUR INTERVAL 20 FEET DATUM IS MEAN SEA LEVEL

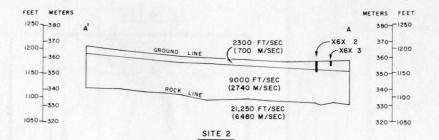
BASE ADAPTED FROM U.S. GEOLOGICAL SURVEY TOPOGRAPHIC QUADRANGLE MAPS

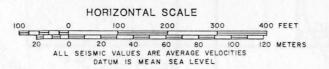
Figure 2 .-- Detailed location maps of seismic surveys (continued)



SITE I

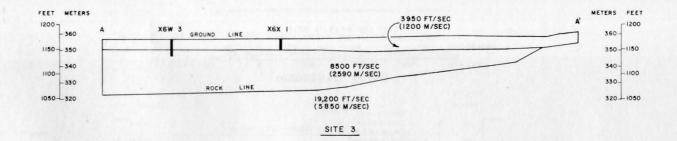


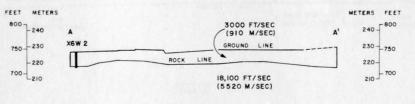




These profiles are the result of a seismic survey made during June 1973 for Vermont Department of Water Resources by Lockwood, Kessler & Bartlett, Inc. The survey was made at selected areas in the upper Winooski Basin to determine the thickness of unconsolidated sediments overlying bedrock. Locations of the seismic lines corresponding to the profiles are shown on figure 2. Low velocity (1500-4000 ft/s or 457-1219 m/s) layers at the top of the profiles are usually interpreted as unconsolidated sediments which may lie above the water table. Medium velocity (4700-5300 ft/s or 1432-1615 m/s) layers are interpreted as till or bedrock.

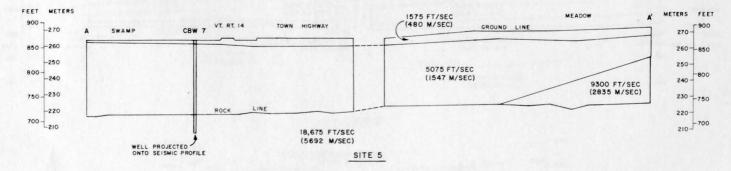
The ground line (assumed) designates the profiles which were not surveyed prior to the set-up of the seismic instrument.





SITE 4

-7-



HORIZONTAL SCALE

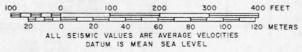
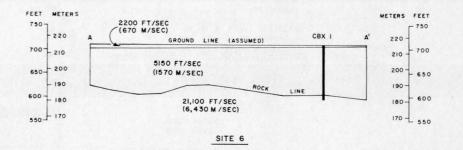
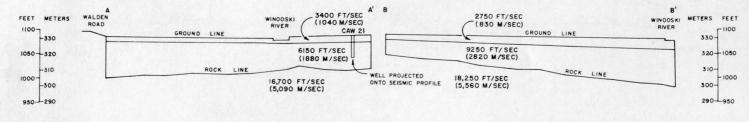
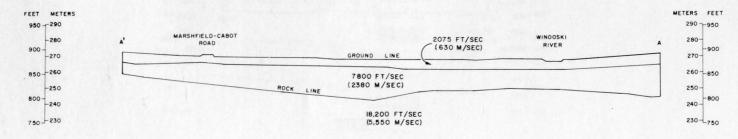


FIGURE 3.--PROFILE SECTIONS OF SEISMIC SURVEYS (CONTINUED)





SITE 7

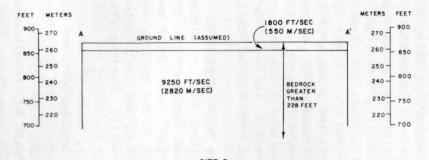


SITE 8

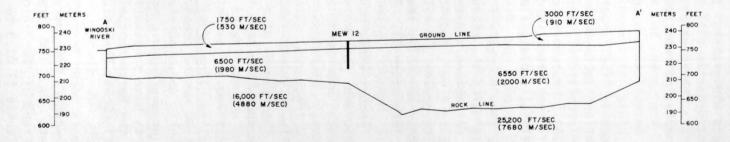
HORIZONTAL SCALE

100 100 200 300 400 FEET 0 चच्चच 1 P 1 20 0 20 100 120 METERS 40 60 80 ALL SEISMIC VALUES ARE AVERAGE VELOCITIES DATUM IS MEAN SEA LEVEL

FIGURE 3.--PROFILE SECTIONS OF SEISMIC SURVEYS (CONTINUED)



SITE 9



SITE IO

HORIZONTAL SCALE

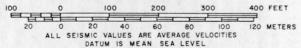


FIGURE 3. -- PROFILE SECTIONS OF SEISMIC SURVEYS (CONTINUED)

-9-

Acknowledgments

The assistance of many individuals and agencies has greatly facilitated the progress of this investigation. The authors wish to express their thanks to the many landowners who granted permission for various tests on their property. The Vermont Department of Health provided access to water-quality data. Special thanks are due those water well drillers who provided information through their well completion reports.

BEDROCK

Bedrock formations in the upper Winooski River basin are primarily metamorphic rocks of Cambrian to Devonian age. Detailed geologic descriptions are provided in the map reports of Doll and others (1961), Cady (1956), Hall (1959), and Konig (1961). Metamorphic rocks, which underlie approximately the western 80 percent of the study area (see plate 1), are compact and altered, and consist of phyllite, schist, slate, quartzite, and slightly metamorphosed limestone. The older Stowe and Missisquoi Formations are separated by an unconformity from the Shaw Mountain Formation and the younger formations above. This unconformity also marks the change between the green argillaceous rocks of Stowe and Missisquoi Formations and the brown-weathering calcareous rocks of the Shaw Mountain, Waits River, and Gile Mountain Formations above.

Granitic intrusive bodies underlie approximately the eastern 20 percent of the study area and occur as small, isolated bodies in the north-central and south-central parts (see plate 1). The large body in the east is the Knox Mountain pluton.

Throughout the report area, most of the individual water users depend on wells drilled into bedrock for water supplies. In the metamorphosed rocks, virtually all pore space has been eliminated between the individual grains which make up the rock. Therefore, in these rocks water moves principally through the joints and fractures. Similarly, in the massive igneous intrusive rocks water movement is in joints and fractures; and their interconnection and saturation determine the yield of a bedrock well. Bedding and schistosity of the metamorphic rocks appear to have little or no effect on well yield.

In areas of thin, unconsolidated deposits, zones of concentrated joints or fractures can be identified by use of aerial photographs and topographic maps. These features are called "lineaments" and appear on maps and photographs as lines or narrow zones of marked topographic or tonal change. Wetter or dryer conditions are frequently noted on the lineaments by tonal contrast. Recent studies in Delaware (Woodruff and others, 1972) have shown that wells drilled to intersect lineaments have substantially higher yields than wells located at random. A study in the Barre-Montpelier area (Hodges and others, 1976a) also suggests a correlation between higher yields and the proximity of wells to lineaments.

Although lineaments occur throughout the area, the frequency of occurrence is greatest in the western 80 percent of the area underlain by metamorphic rocks as shown on plate 1. The predominant trend is northwest--striking at right angles to the trend of the regional structure.

Median yields (table 1) of 5 to 6 gal/min (0.32 to 0.38 L/s) for wells in bedrock in the upper Winooski River basin are generally lower than yields for similar rock types in the Barre-Montpelier area (Hodges and others, 1976a) or the White River Junction area (Hodges and others, 1976b). These lower median

yields may be substantially influenced by the very small number of high volume commercial-industrial water users in the report area. However, a well yield of 5 to 6 gal/min (0.32 to 0.38 L/s) is generally adequate for domestic needs.

The total thickness of unconsolidated deposits which, in the upper Winooski River basin, varies from zero at bedrock outcrops to more than 275 ft (84 m) in Calais (well CBW52), can directly affect both the availability of water and the cost of construction of a water system. Thick deposits of till or fine-grained sediment may retard the movement of water into bedrock fractures, while thick saturated deposits of sand and gravel can readily recharge the bedrock aquifer. Where the unconsolidated deposits are fine arained or have insufficient saturated thickness to yield adequate, dependable quantities of water to fit the user's needs, it is necessary to case through them and drill into bedrock. The length of casing required to case off the unconsolidated deposits can have a significant impact on the total cost of the water system. Estimates of the length of casing needed can be determined from the map showing the thickness of unconsolidated deposits (plate 2) prepared from analysis of well-construction data in table 6, lithologic logs in table 7, and seismic-survey data in figures 2 and 3.

The data in table 1 on median depth of wells drilled in each formation suggest that where massive, generally brittle rock types are present, such as quartzite or granite, adequate yields may be obtained with shallower wells than in areas underlain by incompetent phyllite and schist. Wells were shallowest in the Moretown Member of the Missisquoi Formation, with a median depth of 130 ft (40 m) for 35 wells. The next deeper wells were located in the granitic intrusive areas, followed by wells in the Barton River Member of the Waits River Formation, with the deepest wells being located in the Gile Mountain Formation. The median depth for 19 wells in this last formation is 200 ft (6) m). There appears to be a direct relationship between the massiveness or competence of rock type, the degree of jointing and fracturing, and the extent to which these breaks in the rock remain open to transmit water. It is expected, however, that with depth there is a diminishing probability of increasing well yield because the weight of overlying rock closes the joints and fractures.

Type of well	Number		eld /min)		pth t)	Median* yield/foot
	wells	median	range	mediar	range	[(gal/min)/ft]
Wells finished in bedrock:						
Granite (nhu)	10	5	1-35	150	63-225	0.032
Gile Mountain Formation (Dg) Waits River Formation (Dwb)	19	6	.5-300	200	100-320	.030
Barton River Member Missisquoi Formation	62	6	1-100	174	21-340	.034
Moretown Member (0mm) Domestic wells finished in	35	6	0-50	130	35-500	.038
unconsolidated material	26	18	1-90	58	6-155	.310

Table 1.--Range, median yield, and depth of selected wells by type of material

*Median yield per foot drilled (total depth).

The unconsolidated deposits found in the upper Winooski River basin are a result of several periods of glaciation during Pleistocene time. Moving ice removed soil and rounded and shaped the bedrock surface as the ice flowed over the study area. The ice also deposited a mantle of till over bedrock. Till is an unsorted mixture of rock fragments which range in size from clay to boulders.

Most areas throughout the upper Winooski River basin are covered with till, with moderately to very thick deposits (as much as 275 ft, 84 m) being located near South Cabot, the southeastern section of Marshfield, and substantially filling Carr Brook valley located west of the village of East Calais.

Two types of till have been described in the upper Winooski River basin area by Stewart and MacClintock (1969). Basal till is compact, gray, and commonly fissile, suggesting a subglacial origin. Ablation till is a loose mixture of brown sand, cobbles, and boulders containing minor amounts of silt and clay. Stewart and MacClintock (1969) ascribe the formation of this material to slow settling of supraglacial debris during ice wasting. Water velocities were assumed to be only fast enough to remove clay and silt while leaving the larger particles behind.

Glacial meltwaters also carried, sorted, and deposited rock and soil debris. Streams in channels in and under the melting glacial ice formed long, sinuous deposits of sand and gravel. Deposits of this type are potential aquifers and are found throughout the length of the valley of the Kingsbury Branch and Cooper Brook from Hardwick to Plainfield. Streams also deposited layers of silt, sand, and gravel in some valley areas along the margins of melting ice. Deposits of these types are potential aquifers and are found throughout the upper Winooski River basin (plate 3). The largest is located around Nichols Pond on the Woodbury-Hardwick Town boundary. Large, temporary lakes were formed in most river valleys. Silt- and clay-laden waters entering these lakes deposited fine-grained sediment over the lake bottoms which remain as large flats and terraces in valleys today. The most prominent terraces are located along the lower Kingsbury Branch, Cooper Brook, and the Winooski River from Marshfield to Plainfield. Some terraces are also found along the Middle Branch in the Town of Worcester.

Water Availability in Unconsolidated Deposits

Water fills the pores between grains of unconsolidated deposits below the water table, but not all of it may be available for use. The rate at which water will flow through a deposit (hydraulic conductivity) is a major aquifer characteristic which controls well yield. Generally, coarse-grained and well-sorted deposits have the highest hydraulic conductivities and will provide the highest yields to wells. Silt, clay, and till with low hydraulic conductivity-ities yield little water to wells.

The sustained yield of an aquifer is governed by its rate of recharge and the amount of water held in storage. The amount of water held in storage is related to the type of material within the aquifer, the areal extent of the aquifer, and its saturated thickness. Delineation of aquifers by predominant material type and approximate minimum saturated thickness (plate 3) was obtained from a study of the surficial geology, drillers' records of water wells, U.S. Geological Survey test borings, and seismic information. Estimates of the total thickness of unconsolidated deposits are presented in plate 2, but insufficient data exist to construct a saturated thickness map of the study area. Because aquifers in Vermont are commonly small, the availability of water from this is ultimately dependent on the rate of recharge. This is discussed more fully under the Recharge and Discharge section of this report.

Ti11

Shallow, large-diameter dug wells are used to obtain water from till. Wells in ablation till usually yield more water than wells in basal till because ablation till has a loose, sandy texture and, therefore, has a higher hydraulic conductivity. Because wells in till are usually shallow and located on hillsides where water-table fluctuations are greatest, they are more susceptible to failure or to a reduction in yield during drought than wells in other aquifers.

Fine-grained Deposits

Unconsolidated deposits over much of the valley areas along the main stem of the Winooski River in Marshfield and Cabot and the lower Kingsbury Branch consist of lake bottom deposits composed of fine sand, silt, and clay. These and similar deposits throughout the area are mapped as fine-grained deposits (plate 3). Some are covered with an unsaturated veneer of sand and gravel. The fine-grained deposits have a low potential for ground-water yield because of their low hydraulic conductivity. However, lenses of coarse-grained, intercalated material occur in them at some locations and may yield moderate quantities of water suitable for domestic and some commercial supplies. Although these fine-grained deposits may not readily yield water to wells, they do hold large quantities of water which are available for slow release to supply adjacent streams and unconsolidated aquifers, or underlying bedrock aquifers.

Sand and Gravel

Sand and gravel deposits, where they have sufficient saturated thickness and are readily recharged, offer the greatest ground-water potential in the study area. The most extensive sand and gravel aquifer is located along the Kingsbury Branch between South Woodbury and East Calais (plate 3). This location and other similar areas underlain chiefly by saturated medium sand to gravel deposits with a total thickness of more than 20 ft (6 m) are capable of yielding more than 200 gal/min (13 L/s), sufficient to meet some commercialindustrial or municipal needs. Areas underlain mainly by saturated fine sand to gravel deposits having a total thickness of more than 20 ft (6 m) should yield 50 gal/min (3 L/s) to 200 gal/min (13 L/s), sufficient to meet light industrial and small public supply requirements as at Woodbury well X6W2 (plate 3 and table 3). Recharge from precipitation alone is commonly inadequate to sustain high yield wells in aquifers with small recharge areas. However, induced infiltration from adjacent surface waters can provide additional recharge for such aquifers to sustain high yield wells.

Estimated Yield of Stratified Deposits

Twenty-three test wells were drilled in stratified drift deposits at selected locations to provide data for estimating potential ground-water yields. Seven test wells penetrated water-saturated granular materials of sufficient thickness (at least 17 ft) to have some potential for ground-water development. Because aquifer tests were not made, estimates of ground-water yield at the seven test sites are based on an indirect method (Lohman, 1972, p. 53) using the lithologic description of materials sampled. Estimated values for hydraulic conductivity (table 2) were assigned to each lithologic unit described in table 7. The assigned hydraulic conductivity was multiplied by the saturated thickness of each unit and the products were summed to provide an estimated transmissivity value for each test well site in table 3.

Table 2.--Estimated hydraulic conductivity for various materials

Material	Estimated hydraulic conductivity (feet per day)
Clay]
Till	1
Silt	1
Silt and very fine sand	1
Silt and clay	1
Silt and gravel	2
Fine sand, very fine sand and silt	2
Clayey fine sand to fine grave]	5
Fine sand with clay layers	5
Fine sand, some clay and gravel	10
Fine and medium sand with clay layers	20
Fine sand and medium sand	25
Sandy till	25
Fine sand to fine gravel	30
Medium and coarse sand, clay layers	
A]]uvium	50
Sand and gravel, some clay Medium sand	60
Medium sand	
Medium sand and coarse sand	125
Medium sand, some fine sand to fine gravel	300
Sand and gravel	500
Fine gravel and sand	
Medium and coarse sand, some gravel and silt	700
Fine grave]	800
Medium and coarse sand and gravel	900
Coarse sand to gravel, some fine to medium sand-	900
Coarse sand and cobbles	1000
Gravel	1000
Cobbles and gravel	1000

[Modified from tables by Lohman (1972, p. 53), and Ryder and others (1970, p. 21).] Well yields were computed from the estimated transmissivity values and the maximum allowable drawdown, selected as the difference between the static water level and 1 ft (0.30 m) above the top of the screen, using the Theis non-equilibrium equation (Lohman, 1972, p. 8). The drawdowns were then adjusted for thinning of the aquifer due to dewatering and the effects of partial penetration of the aquifer by the well screen (Cervione and others, 1972, p. 50-55). If the adjusted drawdown exceeded the maximum allowable drawdown, the calculations were repeated until a well yield was obtained where the resultant drawdown did not exceed the specified limit (table 3).

The estimated yield for each site applies strictly to a well 24 in (610 mm) in diameter, 100 percent efficient, that has been pumped continuously for 200 days. It is also assumed that the horizontal hydraulic conductivity is 10 times greater than the vertical hydraulic conductivity and the average storage coefficient is 20 percent.

Hydrogeologic boundaries, although not considered in these calculations, can also affect well yields and resultant drawdowns. The effects of impermeable boundaries (bedrock, till, clay) decrease well yield by increasing drawdown; recharge boundaries, as a result of induced infiltration from surface-water bodies, increase well yield by decreasing drawdown.

The foregoing methods give qualitative estimates for the potential yield of wells in stratified-drift deposits, but systematic exploratory drilling and aquifer testing are necessary for locating large capacity wells.

Table 3.--Estimated transmissivity and yield at selected test-well sites

[Estimated yields apply strictly to a 100 percent efficient, 24-in (610 mm) well that has been pumped continuously for 200 days with the pumping level maintained at 1 foot (0.30 m) or more above the top of the screen. For a full discussion of the methods used and conditions assumed, see the text.]

		Cardinal State State State State State	Construction of the second	Contraction of the second s	Saugerstand and the second sec	and the second se
CAW21	CBW7	MEX1	X6W2	X6W3	X6W4	X9W1
36	84	23	30	20	35	63
34	67	17	27	17	31	63
10	19	5	5	5	10	10
23	42	11	21	11	21	52
23,000	20,000	10,500	5500	8500	24,000	5000
875	1275	225	150	200	875	275
	36 34 10 23 23,000	368434671019234223,00020,000	3684233467171019523421123,00020,00010,500	36842330346717271019552342112123,00020,00010,5005500	368423302034671727171019555234211211123,00020,00010,50055008500	36842330203534671727173110195551023421121112123,00020,00010,5005500850024,000

RECHARGE AND DISCHARGE

Water movement into and out of the project area can be expressed by the following equation:

 $P = R + ET + \Delta S$

Where P = precipitation, R = runoff, ET = evapotranspiration, and ΔS = change in storage. At the Edward F. Knapp Airport, Berlin, near the study area, precipitation averages 34 in (864 mm) per year, with somewhat more during summer than winter.

Of the precipitation each year, evapotranspiration returns an average equivalent of 14 in (360 mm) of water to the atmosphere (Hodges, Butterfield, and Ashley, 1976a). Evapotranspiration is the sum of direct evaporation of surface water, the sublimation of snow, and the transpiration of living organisms. Most evapotranspiration occurs during the spring and summer growing season with the result that ground-water levels decline during this period as trees and plants remove water from the ground and release it to the atmosphere as water vapor. Killing frosts in September or October end the yearly growth cycle, cause transpiration rates to decline, and result in the rise of groundwater levels.

An average of 20 in (508 mm) of water per year leaves the area as runoff. This includes water that runs directly over the land to the streams, and that which seeps into the ground, recharges ground-water bodies, and then discharges to the streams. Ground-water discharge to streams forms a significant proportion of the total streamflow and sustains flow during periods of little or no rainfall or below-freezing temperatures.

Normally, ground-water levels at central Vermont (Hodges, Butterfield, and Ashley, 1976a) have a seasonal high in March or April, coinciding with melting of the snowpack and break-up of ice in the rivers and a seasonal low in September or October at the end of the growing season. This sequence, however, can be modified by excessive rainfall or drought. The change in ground-water storage over the years is probably negligible because decreases in storage during dry years are offset by increases in storage during wet years.

CHEMICAL QUALITY OF WATER

Chemical analyses of 13 ground-water samples (table 4) and 38 surfacewater samples (table 5) were obtained from several governmental agencies for comparison in this study.

Eight of the analyses of water from wells contained one or more constituents that equaled or exceeded the recommended or maximum allowable limits of concentration adopted by the National Academy of Sciences and National Academy of Engineering (1973) for drinking water supplies. Of the eight samples in which the limits were equaled or exceeded, the recommended limit for manganese was exceeded in four samples, the recommended limit for iron was exceeded in two samples, and the maximum allowable limit for nitrate was exceeded in three samples.

Surface water in the upper Winooski River basin was sampled at 23 locations (plate 4) to determine chemical composition. Analyses of these samples are shown on table 5. Surface waters showed a median pH of 7.7 and a median

Table 4.--Chemical analyses of ground water

(Analyses in milligrams per liter except as indicated.)

Local well number	Date	Dissolved iron (Fe)	Dissolved manganese (Mn)	Dissolved sodium (Na)	Alkalinity as CaCO ₃	Dissolved chloride (Cl)	Dissolved nitrate (N)	Dissolved nitrite (N)	Hardness (Ca,Mg)	pH (units)	Color (Platinum- cobalt units)	Turbidity (Jtu)	Total copper (Cu)	Total lead (Pb)	Total zinc (Zn)	Source of data ¹	Type of aquifer ²
Stand	ards ³	0.3*	0.05*			250*	10.0*				15*	5*	1.0*	0.05**	5.0		
САВОТ																	
W-21	7-26-73	0.02	0.09		120	34	18	0.01		7.3	5					WR	U
W-23	6-12-75	.00	.00	4	97	13	.0	.00	120	7.8			0.1	0.00	0.0	н	R(Dg)
	CALAIS																
W-30	7-14-71	.03	.00	2	100	6	1.8	.00	240	7.2	0	0	.0			н	R(Dw)
							м	ARSHFIEL	_D								
W-45	5-14-72	.00	.00	1	60	3	1	.00	66	7.1	0	0	.0	.00	.2	Н	R(Dw)
W-48	1-15-74	.21	.03	8	140	3	.2	.0	140	7.2	0	0	.0	.0	.0	н	R(Dg)
								WOODBURY	'								
W-2	7- 2-73	.00	1.10		110	1.4	.1			6.6	5					WR	U
W-3	7-31-73	.00	.00		120	39	22	.00		7.7	10					WR	U
								ORCESTER	\$								
W-1	8- 3-73	.16	.00		65	44	13	.10		7.5	50					WR	U
W-3	1-26-73	.00	1.60										.2	.00	.5	Н	R (Omm)
W-9	11- 7-71	.05	.00	2	50	0	.0	.0	140	6.7	0	0	.0	.0	.0	н	R (Omm)
W-37	12-29-71	.66	.04	4	42	25	.2	.01	54	6.9	0	4	•0	.05	•4	Н	
W-49	8-19-69	.07	.18	1						7.9	5	0	.06	.00	.018	US	U
W-50	7-11-74	.66	.03	96	68	150	1.8	.0	92	6.4	0	0	.0	.0	.0	Н	R (Omm)

ISource of data: H, Vermont Department of Health; WR, Vermont Department of Water Resources; and US, U.S. Public Health Service.

²Type of aquifer: R, bedrock (letters in parentheses refer to bedrock formations, see Plate 1); U, unconsolidated materials.

³National Academy of Sciences and National Academy of Engineering (1973) drinking water standards limits of concentration: *recommended limit, **maximum allowable.

Table 5.--Chemical analyses of surface water

(Analyses by Vermont Department of Water Resources. Analyses in milligrams per liter except as indicated.)

Map no.	Location and station number ¹	Date	Dissolved calcium (Ca)	Dissolved magnesium (Mg)	Alkalinity as CaCO ₃	Dissolved chloride (C1)	Hardness (Ca,Mg)	pH (Units)	Temperature (°C)	Color (Platinum- cobalt units)	Turbidity (Jtu)	Dissolved oxygen	Total coliform (Most probable number colonies per 100 milliliter)
1	Winooski River (29)	8-31-53 8-24-65	80	12	88	5.5	92	7.8	23.0 18.0	10 20	05	7.8 9.0	1,500
2	Nasmith Brook (30)	9-03-53	74	14	84	3.0	88	8.3	18.5	5	1	9.5	450
3	Nasmith Brook (31)	9-03-53	74	24	86	2.4	98	8.2	18.0	0	0	9.3	250
4	Winooski River (32)	8-31-53 7-15-55 6-06-57	38 38 54	22 2 10	40 42 44	3.6 1.4 6.0	60 40 64	7.7 7.6 7.7	20.0 19.0 15.0	25 30 20	1 14 4	8.4 8.3 9.2	9,500 450,000
5	Marshfield Pond Brook (33)	9-03-53	18	44	24	2.4	62	8.0	21.0	25	0	8.7	4,500
6	Winooski River (34)	8-31-53 6-06-57 8-24-65	38 48	26 22 	40 44	4.2 6.0 5.5	64 70	7.3 7.7 8.0	19.7 14.5 18.0	25 20 35	544	6.5 8.4 8.0	9,500
7	Molly's Falls Brook (35)	9-03-53	88	8	94	5.4	96	8.1	19.0	10	0	8.4	250
8	Winooski River (36)	7-15-55 6-06-57 6-21-62	118 98	22 88	142 40	3.4 5.0 5.0	140	7.7 7.3 7.6	20.5 14.0 16.5	35 30 15	38 6 3	5.2 3.4 7.4	1,100,000
9	Winooski River (37A)	6-21-62				2.0		7.5	16.0	20	3	6.2	80,000
10	Winooski River (37)	8-31-53 6-06-57	106 100	22 26	114 98	6.1 6.0	128 126	6.9 7.5	23.0 13.5	20 20	20 5	.3 5.8	950,000
11	West Hill Pond Brook (38)	9-03-53	88	20	94	3.0	108	8.0	19.0	10	0	8.4	1,500
12	West Hill Pond Brook (39)	9-03-53	78	12	72	3.0	90	7.4	24.0	15	1	7.1	7,500
13	Winooski River (40)	8-31-53 7-15-55 6-06-57 3-13-63 8-24-65	144 208 100	32 44 16 	184 58 100 	18 20 6.0 4.4 8.3	176 252 116 	6.7 5.2 7.5 6.8 7.9	24.0 21.0 12.5 1.2 17.0	15 15 30 10 20	47 60 6 6 14	0 6.6 10.9 7.0	2,500,000 1,100,000
14	Winooski River (41)	8-31-53 7-15-55 6-21-62 3-13-63	140 84 	16 22 	145 130 	3.6 .7 .8 2.8	156 106 	7.8 8.0 8.1 8.1	19.5 19.0 14.0 .5	5 10 10 10	0 6 1 6	7.5 8.7 10.2 13.0	950 2,500 2,400
15	Kingsbury Branch (7-4)	9-04-53	98	24	108	7.3	122	7.9	20.5	15	0	6.8	2,500
16	Kingsbury Branch (7-5)	9-04-53	92	22	100	7.3	114	8.2	21.0	10	0	7.9	9,500
17	Kingsbury Branch (7–6)	9-04-53	96	16	96	6.7	112	7.9	16.0	0	0	8.0	9,500
18	Pekin Brook (7A-1)	9-04-53	102	22	118	2.4	124	7.8	22.0	15	5	6.7	2,500
19	Curtis Pond Brook (7A-2)	9-04-53	80	24	86	3.0	104	7.8	22.0	25	0	6.7	750
20	Sugar Brook (7A-3)	9-04-53	106	14	116	3.0	120	7.6	18.5	0	0	6.6	9,500
21	Pekin Brook (7A-4)	9-04-53	72	10	70	3.0	82	8.1	21.0	0	0	8.3	9,500
22 23	North Branch (5-8) North Branch (5-9)	7-11-55 7-11-55	16 16	8	18 18	1.0	24 18	7.0 7.0	22.0 25.5	20 35	4	6.8 8.4	4,000 75,000

¹Station numbers are site designations used by the Vermont Department of Water Resources.

hardness of 105 mg/L (milligrams per liter) as CaCO₃. This compares with a median pH of 7.2 and hardness of 120 mg/L for ground-water samples. Surface water is generally more alkaline than ground water. The most alkaline of the surface-water samples came from three streams originating principally from the area of the Knox Mountain pluton in the southeast corner of the study area.

Hardness

Hardness is a term applied to the soap-neutralizing power of water. Hardness is largely determined by the concentration of calcium and magnesium, and is expressed as calcium carbonate (CaCO₃). The following ranges have been used in various U.S. Geological Survey reports to classify hardness:

Hardness as CaCO ₃ (mg/L)	Descriptive rating
0 - 60	Soft
61 - 120	Moderately hard
121 - 180	Hard
181 or more	Very hard

While soft and moderately hard water can generally be used without treatment except for some industrial purposes, hard water frequently requires treatment (softening) for use in laundries, some industrial and most domestic uses. Very hard water requires softening for most purposes to make the water usable and to prevent damage to water-supply and water-using equipment.

Median hardness of the six ground-water analyses in the upper Winooski River basin was 120 mg/L, ranking the water as moderately hard to hard. Therefore, treatment of some water may be required.

The degree of hardness shown by these ground-water samples generally reflects the regional trends found in surface-water samples, wherein surface waters originating from calcareous rocks (Gile Mountain and Waits River Formations) show the highest levels of hardness (106 to 252 mg/L). Surface waters flowing from the area of the Knox Mountain granite show lower levels of hardness (88 to 96 mg/L) and streams underlain by quartzite and slate (Moretown Member of the Missisquoi Formation and Stowe Formation) show the lowest levels of hardness (18 to 24 mg/L).

Iron and Manganese

Iron and manganese are minor constituents of water, but excessive concentrations, particularly of manganese, may have harmful effects on health (Yase, 1972). The National Academy of Science and National Academy of Engineering (1973) recommended a maximum limit of 0.3 mg/L of iron and 0.05 mg/L of manganese in drinking water supplies. The median iron content for all wells sampled (table 4) was 0.03 mg/L, although samples from two rock wells were more than double the recommended limit. Iron found in the water may be from many sources in the upper Winooski River basin. Most bedrock formations in the report area contain iron-bearing minerals. Magnetite is listed as a common accessory mineral by Cady (1956) and Konig (1961), and was commercially extracted from sand deposits near East Calais during the 1800's (Konig, 1961). Iron in ground water may increase as a result of reducing conditions produced by decaying organic matter in aquifers or derived from industrial waste and dumps (landfills). Median manganese content for wells sampled was 0.03 mg/L, which approaches the limit of 0.05 mg/L recommended by the National Academy of Science and National Academy of Engineering. However, three domestic wells and one public water-supply well sampled showed excessive manganese. Manganese is not a major part of the mineral composition of any rock types identified by Cady (1956) or Konig (1961) in the upper Winooski River basin. It may be derived from minerals in which it is a minor constituent, or from industrial wastes and dumps. The lack of an obvious source of the manganese suggests that only an in-depth study could determine the factors contributing to excessive levels shown in table 4.

SUMMARY

Ground water can be obtained in quantities suitable to sustain singlefamily domestic and farm supplies from wells drilled in bedrock nearly everywhere in the upper Winooski River basin. The median yield for 126 domestic supply wells in four different bedrock formations is between 5 and 6 gal/min (0.32 to 0.38 L/s). The median depths for wells in the four different formations range from 130 to 200 ft (40 to 61 m). In the area studied, bedrock well yields are more dependent on rock fracture than rock type; generally, wells located in fracture zones have the greatest yields. Some fracture zones appear as linear features (lineaments) on aerial photographs or topographic maps (plate 1). Lineaments may be used for well-site selection to enhance the probability of obtaining higher-than-average well yields.

Water-saturated sand and gravel aquifers capable of yielding up to 1,000 gal/min (63 L/s) to individual wells, sufficient to sustain commercial, industrial, or municipal supplies, are found in valley areas of all five towns in the study area. These potential sources of large supplies are found throughout the length of the valley of the Kingsbury Branch and Cooper Brook from Hardwick to Plainfield. Water availability and location of sand and gravel aquifers in the valley areas are described on plate 3.

Of the 23 test wells drilled for this study, 7 penetrated water-saturated sand and gravel of sufficient thickness to have potential for ground-water development. Well yields for these seven sites, calculated from geohydrologic data from the test wells, ranged from 150 to 1,275 gal/min (9.5 to 80 L/s).

Chemical analyses of 13 ground-water samples and 38 surface-water samples were used to evaluate water-quality conditions in the basin. Eight of the 13 analyses of ground water revealed one or more constituents that equaled or exceeded the maximum allowable limits for iron, manganese, and nitrate as nitrogen in drinking water. Iron and manganese occur naturally, but nitrate is an indication of pollution from human or animal wastes or from fertilizer. Surface water and ground water were found to be moderately hard to hard. Hardness is derived from carbonate minerals in the bedrock and overlying unconsolidated deposits, and can be reduced through treatment.

Ground-water resources in the basin are of adequate quantity and quality to meet foreseeable needs. Although resources capable of sustaining municipal supplies occur in the basin, they are located only in some valley areas. Test drilling and exploration within these aquifers would be a necessary forerunner of any development. Except for localized water-pollution problems and some instances of excessive levels of iron, manganese, or hardness, which can be controlled with treatment, the chemical quality of ground water is suitable for drinking. LOCAL WELL NUMBER: LETTER PREFIX INDICATES--A, U.S. GEOLOGICAL SURVEY AUGER BORING; B, BRIDGE BORING; R, ROADWAY BORING; W, WELL OR TEST WELL; X, MISCELLANEOUS TEST BORING.

LATITUDE-LONGITUDE: NUMBER FOLLOWING DECIMAL POINT IS A SEQUENTIAL NUMBER FOR WELLS OR BORINGS IN A 1-SECOND GRID.

ALTITUDE OF LAND-SURFACE DATUM: ALTITUDES ARE EXPRESSED IN FEET ABOVE MEAN SEA LEVEL.

METHOD DRILLED: A, AIR-ROTARY; B, BORED OR AUGERED; C, CABLE TOOL; D, DUG; H, HYDRAULIC-ROTARY; J, JETTED; P, AIR-PERCUSSION; R, REVERSE-ROTARY; T, TRENCHED; V, DRIVEN; W, DRIVE-WASH.

WELL FINISH: C, POROUS CONCRETE; F, GRAVEL WALL WITH PERFORATED OR SLOTTED CASING; G, GRAVEL WALL WITH COMMERCIAL SCREEN; H, HORIZONTAL GALLERY OR COLLECTOR; O, OPEN END; P, PERFORATED OR SLOTTED CASING; S, SCREEN; T, SAND POINT; W, WALLED OR SHORED; X, OPEN HOLE IN AQUIFER (GENERALLY CASED TO AQUIFER).

WELL DEPTH: DEPTH OF FINISHED WELL, IN FEET BELOW LAND SURFACE.

WELL USE: A, ANODE; D, DRAINAGE; G, SEISMIC HOLE; H, HEAT RESERVOIR; O, OBSERVATION; P, OIL OR GAS; R, RECHARGE; T, TEST; U, UNUSED; W, WATER WITHDRAWAL; X, WASTE DISPOSAL; Z, DESTROYED.

WATER-BEARING MATERIAL: PRINCIPAL WATER-BEARING ZONE.

	ADJECTIVE (FIRST CHARACTER)	LI	THOLOGY (SECOND CHARACTER)
1	VERY FINE GRAINED	A	ALLUVIUM
2	FINE GRAINED	В	SEDIMENTARY ROCK,
3	MEDIUM GRAINED		UNCLASSIFIED
4	COARSE GRAINED	С	CONGLOMERATE
5	VERY COARSE GRAINED	D	DOLOMITE
6	CLAYEY	E	GYPSUM OR ANHYDRITE
7	SILTY	F	SHALE
8	SANDY	G	GRAVEL
9	GRAVELLY	Н	IGNEOUS, GRANULAR
0	CAVERNOUS		(GABBRO, GRANITE, ETC.)
A	ARGILLACEOUS	I	IGNEOUS, APHANITIC OR
В	BOULDERY		GLASSY (BASALT, ETC.)
C	CALCAREOUS	J	IGNEOUS, UNCONSOLIDATED
D	DENSE		(TUFF, VOLCANIC ASH)
E	CONCRETIONARY	K	SAPROLITE
F	IRON STAINED OR IRON CEMENTED	L	LIMESTONE
G	GRANULAR	м	MARL OR SHELL MARL
н	HARD	N	METAMORPHIC, COARSE
1	INTERBEDDED		GRAINED (GNEISS, MARBLE,
J	JOINTED OR FRACTURED		QUARTZITE)
ĸ	COLUMNAR	0	METAMORPHIC, FINE GRAINED
L	LAMINATED OR BANDED		(SCHIST, SLATE)
M	MASSIVE	Р	CLAY
N	NONCALCAREOUS	Q	SILT OR LOESS
0	ORGANIC	R	SAND AND GRAVEL
P	POORLY SORTED	S	SAND
Q	CHERTY OR SILICEOUS	Т	TILL
R	REDBED	U	UNCONSOLIDATED SEDIMENT
S	SOFT	v	SANDSTONE
Т	"SALT AND PEPPER"	W	SILTSTONE
U	UNCONSOLIDATED	X	SILTY SAND
V	SEMICONSOLIDATED	Y	CLAYEY GRAVEL
W	WELL SORTED	Z	OTHER
X	CROSS BEDDED		
Y	SHALY OR SLATY		
Z	WEATHERED		

WATER LEVEL: LEVELS ARE GIVEN IN FEET BELOW LAND SURFACE; "+" INDICATES WATER LEVEL ABOVE LAND SURFACE; "F" INDICATES FLOWING WELL.

WATER USE: A, AIR CONDITIONING; B, BOTTLING; C, COMMERCIAL; D, DEWATERING; E, POWER GENERATION; F, FIRE PROTECTION; H, DOMESTIC; I, IRRIGATION; M, MEDICINAL; N, INDUSTRIAL (INCLUDES MINING); P, PUBLIC SUPPLY; R, RECREATION; S, STOCK; T, INSTITUTIONAL; U, UNUSED; V, REPRESSURIZATION; W, RECHARGE; X, DESALINATION--PUBLIC SUPPLIES; Y, DESALINATION--OTHER SUPPLIES.

"PUMPAGE/YIELD: IN GALLONS PER MINUTE (GAL/MIN).

PUMPAGE/DRAWDOWN: THE DIFFERENCE BETWEEN STATIC WATER LEVEL AND PUMPING LEVEL.

PUMPAGE/TIME: THE FOLLOWING CODES ARE USED FOR PUMPING PERIODS OF LESS THAN 1 HOUR: A, THROUGH 15 MINUTES; B, 16 TO 30 MINUTES; C, 31 TO 45 MINUTES; D, 46 TO 59 MINUTES.

LOG: D. DRILLER'S LOG; E. ELECTRIC LOG; G. GEOLOGIST'S LOG (LOG AVAILABLE IN TABLE 7).

QW: TYPE OF CHEMICAL ANALYSIS AVAILABLE IN TABLE 4, C, COMPLETE; J, CONDUCTANCE AND CHLORIDE; K, CONDUCTANCE; L, CHLORIDE; M, MULTIPLE (INCLUDES ONE COMPLETE AND ONE OR MORE PARTIAL); P, PARTIAL.

TABLE 6 .- - DESCRIPTION OF SELECTED WELLS. TEST WELLS. AND BORINGS -- CONTINUED

							WE	LL		FEET	WATER-		WATER		PL	MPAG	E		
LOC WE NUM		LATITUDE- LONGITUDE	ALTI- TUDE OF LSD (FT)	OWNER OR USER	YEAR/ METHOD DRILLE	(IN)	I ISH	IDEPTHI I I I (FT)I	USE	TO	BEARING		IMEAS-I				1	LOG	0.
*****	12345	442437N0721815.1 442402N0721850.1 442402N0721850.2 442402N0721851.1 442337N0721954.1	1210 970 950 1020 940	CABOT TOWN CABOT CREAMERY CABOT CREAMERY CABOT CREAMERY MORSE CECIL	1949 (1938 (1960 - 1960 - 1965 (2 3 8 8 - 6	X X X X X W	225 320 300 200 6	**C**	15 5	QL QL 25	F F 		IZCZU	150 15 15 5	50		- - - D*	
*****	6 7 8 9 10	442336N0721950.1 442214N0722019.1 442327N0721955.1 442335N0721957.1 442330N0721956.1	930 910 950 940 945	BRIMBLECOMBE R BEAN GERALD BICKFORD A E LAMPHERE HAROLD MORSE CECIL	1973 1972 1973	6 6 6 7 6 6 6 6	× 0 × × ×	124 61 255 122 209	*****	 13 35 6	QL G 	12 21 20 8	8-73 11-72 4-68	IIIIO	3 6 2 3 12			- D - D :	
*****	11 12 13 14 15	442325N0722003.1 442334N0721927.1 442351N0721857.1 442401N0721342.1 442453N0721358.1	940 1110 1180 1620 1660	CARPENTER HAZEL PISTONE LEON CABOT CREAMERY SVEL GEORGE INCARNATION D	1970 1971	6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	x	169 262 342 245 185	*****	6 17 20 6 31		23 40 F 135 20	4-68 10-70 8-70 6-71 1-72	IIZII	25 2 33 2 6			- 0 0 0 0 0 0	
*****	16 17 18 19 20	442638N0721824.1 442615N0721810.1 442455N0722055.1 442427N0722004.1 442349N0721713.1	1460 1360 1280 1360 1720	PIKE PHILLIP SEARLES ROBERT CHATEN E W BOTHFIELD T BURMINGHAM W E	1972	рр 6 6 6 6 6 6 6 6	x	115 164 202 223 162	11111	5 6 9 8 9	=	15 8	2-71	TITI	10 1 3 12			D D O O O	
3 3 3	53 53 51	442435N0721848.1 442312N0722005.1 442420N0721800.1	1080 920 1280	US GEOL SURVEY MIDDLETON DAVID CABOT TOWN		w 1 p 6 p 6	T X X ALAIS	36 121 295		55 8	R 	2 6 F	7-73 10-72 6-68	DID	25 7 300		2 1 	DG3 D D	с р - Р
*****	1 2 3 4 5	441830N0722630.1 441912N0722650.1 441836N0722648.1 441834N0722653.1 441847N0722652.1	730 770 730 750 780	BARTLETT OTTO COTEY PAUL EATON ELIZABETH FARNSWORTH EARL ABBOTT MILTON	1966 1950 1964	CP 6666	S X X 0 0	150 150 128 25 30	33333		=	10 10	-49 -63	IIIII	83 5 4				
	6 7 8 9 10	442042N0723057.1 442238N0722514.1 442645N0721904.1 441912N0722710.1 441912N0722710.2	1270 880 780 860 860	LANE BRADFORD US GEOL SURVEY CAMPBELL C CALAIS ELEM SCH CALAIS ELEM SCH	1973 1970 1970	Р 6 W 1 Р 6 Р 6 Р 6	X T X X X	37 95 140 190 205	*****	* 1 39 20 16	R 	18 16 30 15 7	4-67 7-73 7-70 5-70 5-70	SUHTT	25 12 6 12 4		2 B 2 2	DG: D D D	
****	11 12 13 14 15	442012N0722647.1 442018N0722647.1 44204N0722737.1 442053N0722735.1 442243N0722415.1	780 760 740 800 1360	COAN NILES COFFRIN JOHN COLE VIRGINIA MURELL FORREST FRANKLIN DAVID	1972 1968 1970	P 6 6 6 P 6 6 7	x x x x x x x	182 130 250 210 218	*****	150 100 49 65 50	=	30 12 0 10 4	10-68 5-72 10-68 6-70 7-70	TITIT	20 20 100 4 1		1 1 2 8 8	D* D D D D	
	16 17 18 19 20	442244N0722510.1 442328N0722640.1 442012N0722914.1 442129N0722837.1 442155N0722843.1	890 1050 1280 940. 900	ELURED OWEN DAILEY & BRYANT HOLMQUEST STOWELL JUNE CHESAUX OLIVIER	1970 1973 1971	P 6676	0 x x x x x	81 260 340 255 160	33333	30 21 6 70	G 	5 32 12 10 40	10-70 7-70 8-73 11-71 11-69	TITI	60 1.8 8 2 2		1 1 8 1 1	00000	
****	21 22 23 24 25	442144N0722929.1 442058N0722951.1 442100N0722958.1 442003N0722900.1 442003N0722900.2	1190 1260 1240 1200 1200	OHMEN PAUL SCHOFF CHARLES BUELL HAROLD L RUDIN ANDREW COPELAND W	1973 1973 1968	P 6 6 7 6 6	x x x x x x	220 116 165 21 200		6 10 6 3	=	45 6 20 5	5-68 6-73 9-73 11-68	TITIT	1.5 6 3 12 1		 1 1 1		
* * * * *	26 27 28 29 30	442003N0722900.3 442213N0722953.1 442230N0722944.1 442245N0722930.1 442258N0722922.1	1200 1300 1200 1290 1280	WINSTON JOHN PURCHASE R MAPLE CORNER ST BETZ EDWARD MORSE EVA	1973	- 6 6 7 6 7 6 6	* * * * * * *	123 130 100 160 100	****	24634	=	10 15 7 10 15	6-73 11-73 7-73 7-70 11-69	TITI	6 8 7 6 5		1 2 1 1 2	000000	P
	31 32 33 34 35	442310N0722933.1 441943N0723001.1 441949N0723013.1 441948N0723017.1 442027N0723037.1	1080 1080 1080	DICKENSON E SLAYTON ELGIN SUCHOMEL FRANK ADAMENT MUS SCH SCOTT ROBERT	1969 1973 1970	р 6 6 6 6 6 6 6 7	x x x x x x x	250 70 125 62 160	333333	40 9 4 4 8	=	2 9 10 10	7-70 10-69 6-73 9-71	IIIFI	3 30 8 9 2		1 A 2 1 1	D D D D D	
****	36 37 38 39 40	442039N0723105.1 442126N0723115.1 442202N0723047.1 442242N0723023.1 442104N0723017.1	1550 1480 1390	LANE CLAIR CHERRINGTON J MCBRIDE ALBERT CHERRINGTON J LENO PHILIP	1972 1974 1974	р 6 р 6 р 6 р 6 р 6	* * * * * *	90 75 100 55 142	11111	10 7 6 3 0	=	15 12 12	6-73 5-72 3-74	IIIII	10 7 8 50 8		8 1 8 8	00000	
*****	41 42 43 44 45	442104N0723012.1 442246N0723008.1 442231N0723003.1 442113N0723123.1 442310N0722513.1	1360 1220	LAFOUNTAIN C BROWNELL J DEFORGE DONALD LEVIN HERBERT CALAIS TOWN	1968 1973 1971	р 6 Р 6 Р 7 Р 7 Р 7 6	X X X X 0	130 127 205 230 76	33333	6 4 100 9	 R	10 8 10	2-68 7-73 7-71	ITTTT	2 8 15 5 12		2 1 8 1	0 0 0 0	
*****	46 47 48 49 50	442151N0722428.1 442322N0722647.1 442655N0722407.1 442651N0722106.1 442134N0723130.1	1300 1020 1170 900 1650	STUKE HERBERT L BATES WILLIAM SMITH ROGER LEONARD RUDOLPH RUSSELL WILLIAM	1975 1973 1969	P 6 P 6 P 7 P 6	x x x x x x x	192 198 182 210 190	33333	6 7 4 19 9		20 35 14	9-73 6-69 7-74	IIIII	60 2 6 3 4		1 1 2 8	0 0 0 0 0	

LOG	CAL		ALTI-		YEAR	,	WEL	L		FEET	WATER-		WATER		PL	MPAG	E		
WE	HER R	LATITUDE- LONGITUDE	TUDE OF LSD (FT)	OWNER OR USER	METHO	D	DIAM-IFIN-I ETER IISH I (IN) I I			TO	BEARING MATERIAL		IDATE IMEAS- IURED	1	YIELD	P	1	LOG	0.4
						CA	ALAIS CUNT	INUED											
****	51 52 1 2 3	442235N0722947.1 442217N0722620.1 441905N0722626.1 442215N0722925.1 442215N0723006.1	1220 900 720 1150 1220	SMITH GERALD D SILBERMAN R US GEOL SURVEY US GEOL SURVEY US GEOL SURVEY	1974 1972 1973 1974 1974	Q I X X X	6 X 9 X 2 X 2 X 2 X	149 275 115 15 31	W Z T T T	8 	=			TIDDD	8 		1 	D D DG ^x DG ^x DG ^x	-
							MARSHFIELD												
****	1 2 3 4 5	441824N0722401.1 441919N0722254.1 441952N0722352.1 442001N0722409.1 441948N0722238.1	880 810 1340 1325 790	HAWES J S KIMBALL CHANDLR PITKIN BELMONT PITKIN RONALD ROBERTS RAYMOND	1958 1946 1961 1964	0-090	8 X 6 X 5 X 8 X	194 210 105 140 150	* * * * *	5 100 10 6 98		15 25 25		COLLI	15 35 2 1.5				
*****	6 7 8 11 12	441945N0722240.1 441818N0722402.1 441633N0722213.1 441719N0722434.1 441802N0722345.1	790 880 1330 755 870	CHURCH R C WELCH JOHN WILLARD NELSON DRIVE IN US GEOL SURVEY	1965 1961 1964 1972 1973	A D I D W	6 0 6 X 6 X 6 X 1 T	87 430 120 350 47	33337	20 10	HH R	20 5 3	-65 7-73	TITUD	20 1.5 20 1.5		 2	D× - D DG×	
	13 14 15 16 17	442218N0722300.1 442205N0722245.1 442105N0722326.1 442100N0722223.1 442015N07222447.1	1460 1320 1260 1280 1400	ENNIS LEELAND OATLEY ROBERT BROWN LEON HEALEY JERI MOULTON WILLIAM	1971 1971 1968 1966 1973		7 X 7 X 6 X 6 X 6 X	175 250 125 100 148	*****	3 0 4 15 3		30 28 	3-71 4-71 5-68	TITI	12 1 30 3 2		2 1 1 8 1		
*****	18 19 20 21 22	442020N0722423.1 442016N0722405.1 442009N0722410.1 442015N0722226.1 441950N0722224.1	1300 1260 1320 780 800	CODLING ROBERT HIGGS ALAN PITKIN ROYCE BRIMBLECOMBE S BURNHAM	1969 1970 1968 1969 1970		6 X 7 X 6 X 6 X 6 X	100 195 300 114 100		15 46 8 10 69	=	10 50 50 8 15	8-69 11-70 7-68 10-69 10-70	TITI	1.5 25 0.5 		1 		
*****	23 24 25 26 27	441954N0722239.1 441901N0722525.1 441842N0722522.1 442050N0721807.1 442058N0721807.1	800 1230 1180 1480 1460	POWERS KENNETH WHITCOMB MAHLON FOWLER FRED CHAMBERLAIN C CHAMBERLAIN B	1969 1972 1969 1969 1969		6 X 6 X 6 X 6 X	100 175 220 225 180		10 8 20 30	=	10 15 20 12	6-69 9-72 7-69 6-69	IIIII	30 5 6 1 5		1 2 		
*****	28 29 30 31 32	442016N0721715.1 441936N0721725.1 441749N0722505.1 441742N0722511.1 441750N0722422.1	1620 1750 1060 1000 940	MORSE CECIL CAMPBELL R NUNZIOTO R LOSO LARRY JOHNSON MARTIN	1972 1970 1973 1968 1973		6 0 6 X 7 X 7 X 6 X	41 150 315 255 220	*****	25 17 40 18	R 	40 F 20 15	-70 8-73 5-68 11-73	TITI	4 1 7 1 4		1 8 1 2 2	00000	
****	33 34 35 36 37	441753N0722349.1 441800N0722317.1 441800N0722304.1 441750N0722303.1 441747N0722246.1	800 920 960 1000 1040	TWINFIELD H S SCHROTH RICHARD BLACKBURN GREGG JACOBSEN ERLEND PEARSON JAMES	1969 1968 1968 1969 1972	22220	6 X 6 X 6 0 6 0 6 X	450 142 63 81 151		158 60 100	R R	 15 40	 10-68 9-69	- I I I I	15 5 4 20 5		2 2 2 1	D D D N D N D	
	38 39 40 41 42	441705N0722236.1 441655N0722204.1 441637N0722216.1 441547N0722215.1 441740N0722350.1	1200 1380 1320 1410 760	DAVIS NEIL HORTON THORSTEN NELSON PAUL BOISSE HENRY A BROWN STANLEY	1973 1968 1968 1969 1970		6 0 6 X 6 0 6 0 6 X	54 63 32 54 275		15 87	G G 2G	6 F	6-73 7-70	ITITO	8 4 9 20 6		2 2 1	D × D × D × D × D	
*****	43 44 45 46 48	441710N0722452.1 441633N0722430.1 442142N0722210.1 441615N0722201.1 441857N0722527.1	760 1180 1260 1420 1200	ORTON EDWIN L BATES EVELYN LINDNER DANIEL FRANKS ROBERT RABIN JULES	1970 1968 1970	00101	6 X 6 X 6 X 6 X	160 160 125 225 165	****	70 8 198		1 6 30	7-70 7-68 4-70	IIIII	6 4 30			00-00	
×	1	442126N0721950.1	1060	US GEOL SURVEY	1974	w	2 X	44	T					U				DG×	-
	1 2 3 4 5	442402N0722452.1 442456N0722524.1 442602N0722746.1 442620N0722458.1 442620N0722458.1	935 1040 1175 1260 1480	BEAUCHAMP L C US GEOL SURVEY US GEOL SURVEY US GEOL SURVEY BILLINGHAM D	1941 1973 1973 1974 1968	DEEEO	48 W 1 T 1 T 1 T 6 X	12 30 35 25 180	WTTTW		3R R R	7 5 3 3	7-73 7-73 6-74	ICCCI	5 90 30 10 30	3	B 2 3 	DG× DG× DG× D	
	6 7 8 9 10	442421N0722439.1 442420N0722614.1 442503N0722736.1 44261RN0722424.1 442625N0722456.1	1140 1180 1200 1360 1180	TUCKER ELWIN OSIER PICHARD VOZZELLA M LENARD ROY A COOKSON & HATCH	1968 1973 1970 1973 1968	P 1 P 1 P	6 0 6 X 6 X 7 X	51 325 160 90 150		 10 18 21	R 	14 5 12 15	4-68 8-70 10-73 5-68	TITI	15 15 4 10 10		1 1 1	D # - D D D	
	11 12 13 14 15	442800N0722429.1 442830N0722345.1 442749N0722418.1 442630N0722644.1 442933N0722530.1	1140 920 1480 1260 1440	SELIGA D FLETCHER HOWARD VT FISH & GAME GALLAGHER JOHN WILLIAMS G	1973 1973 1975 1969 1969	0 1 0 0 0	6 0 6 0 6 X 6 X 6 X	155 48 185 130 247		 7 30 24	R R 	20 90	10-73 5-75	TIHII	30 15 15 4 30		1 1 	D # D # D D D	
*****	16 17 18 19 20	442933N0722521.1 442858N0722621.1 442737N0722504.1 442700N0722500.1 442523N0722135.1	1440 1480 1220 1240 1370	WILLIAMS JOE COURCHAINE F POTVIAN RRADLEY MCCOY ALICE D RODWIN LLOYD	1974 1973 1973 1973 1974	P . D . D	6 X 6 X 7 X 6 X	130 101 180 185 173		50 17 36 20 24		18 15 16	9-73 6-73 8-74	IIIII	10 4 6 15 10		1 1 	D D D D D	

TABLE 6 .-- DESCRIPTION OF SELECTED WELLS. TEST WELLS. AND BORINGS -- CONTINUED

	~~~				VEAD		WEL	L			WATER-		WATER		P	UMPAG	E		
	CAL ELL MEER	LATITUDE - LONGITUDE	ALTI- TUDE OF LSD (FT)	OWNER OR USER	METHOD DRILLE	DIA D ETE	R IISH I		USE	TO	BEARING		IDATE I IMEAS-I IURED I	1			1	LOG	QW
			(11)				JRY CON			RUCK			TORED		(GPM)		I (HR)		
* * 5 5 4	21 22 23 1 2	442611N0722530.1 442610N0722540.1 442500N0722448.1 442500N0722742.1 442710N0722516.1	1340 1350 1080 1175 1190	SILK MERTON RUDIN ANDY LAMPHERE C US GEOL SURVEY US GEOL SURVEY	1973	Р - Р ¥ ¥	6 X 6 X 6 X 2 X 2 X	247 182 230 20 21	****	15 4 7 	=	12	10-73	TITOO	1 4 3 		 	D D DG× DG×	
x	3	442710N0722516.2	1190	US GEOL SURVEY	1973	w	s x	8	T					U				DG×	-
						,	ORCESTER				2.59								
* * * * *	1 2 3 4 5	442205N0723307.1 442230N0723258.1 442102N0723331.1 442103N0723340.1 442131N0723251.1	750 740 760 710 1110	US GEOL SURVEY US GEOL SURVEY LAMOUNTAIN W JR MALLERY JOHN DODGE RICHIE	1973 1973 1970 1970 1971	* * ₽ ₽ ₽	1 T 1 T 6 X 6 X 6 X	63 135 201 242 125		136 12 35 5	R 4R 	+2 19 12 8 	8-73 6-73 5-70 9-70	DDIII	45 1 6 1 5		2 9 1 1	DG# DG# D D D	P . P
*****	6 7 8 9 10	442156N0723231.1 442154N0723230.1 442159N0723232.1 442206N0723233.1 442147N0723350.1		MARTIN RUDOLPH DODGE HARRY WELCH SCOTT PERRY LAWRENCE STEFFEN OTTO H	1967 1971 1968 1968 1969		6 X 6 X 6 X 6 X	156 215 126 101 500		14 40 12 8 39		4 40 3 10	3-67 10-71 2-68 9-68	ITITI	4 2 1.5 3 Z		3  2 1 2		
*****	11 12 13 14 15	442147N0723350.2 442142N0723340.1 442227N0723308.1 442231N0723314.1 442232N0723233.1	750 690 790 780 780	STEFFEN OTTO H NELSON JERRY RICHARDSON W MAXHAM SUPPLY DODGE SHELDON S	1969 1974 1971 1973 1968	P P P P P	6 X 6 0 6 X 6 X 6 0	295 61 182 121 102		65 16 45	R	 7 24 F 4	 10-74 3-71 3-73 5-68	ITIZI	Z 50 12 30 20		2 1 1 1 4	D D D D D N D N	
	16 17 18 19 20	442232N0723233.2 442237N0723234.1 442241N0723235.1 442300N0723138.1 442317N0723219.1	760 770	DAY DAVID HEARTHSIDE ENT. MAXHAM DAVID IDOL BEAUREGARD R G	1974 1972 1969 1968 1968	0 0 0 0 0	6 X 6 X 6 X 6 X	149 202 162 190 202	*****	125 85 12 95 10		15 4 0 15 4	6-74 9-72 1-69 6-68 8-68	TITI	6 20 6 10 6		1 1 1 2 2		
	21 22 23 24 25	442403N0723138.1 442403N0723138.2 442405N0723140.1 442408N0723138.1 442552N0723207.1		WILDER DARWIN WILDER WAYNE WILDER RAYMOND LONSKE RONALD DODGE HEATH	1974 1970	0 0 0 0 0	6 X 6 X 6 X 6 X	122 90 142 145 115	*****	2 10 2 8 30	Ē	F 6 10	8-70  8-70 7-70	ITTII	20 25 3 20 25		1 1 2 8	- 0000	
*****	26 27 28 29 30	442418N0723309.1 442359N0723309.1 442257N0723312.1 442250N0723355.1 442305N0723406.1	840 820 780 900 940	NOVAK JOHN C LEE NEWTON JR RICHARDSON W HODGES ARTHUR L MORSE ROBERT	1971 1971 1967 1970 1968	P P P P	6 X 6 X 6 X 6 X	134 141 35 89 82	*****	45 4 7 2 16		50 35 12 6	6-71 10-67 11-70 9-68	IIIII	20 20 9 6 50		1 1 2 1 1		
	31 32 33 34 35	442529N0723348.1 442522N0723409.1 442358N0723423.1 442319N0723423.1 442322N0723423.1	1240 1400 1240 1080 1100	WINTER BARRETT JAMES COURCHAINE P SWEETSER CLYDE HOVEY WILLARD	1972 1968 1972 1973 1973	P P P P P	6 X X X X X X X X X X X X X X X X X X X	225 100 61 180 215	*****	50 23 20 17 15		15 F 20 70	7-68 11-72 8-73 8-73	TITI	1.5 7 4 2 20		B 2 1 1 1		
	36 37 38 39 40	442340N0723515.1 442337N0723512.1 442407N0723531.1 442309N0723418.1 442254N0723427.1	1160 1130 1420 950 1170	SMITH FAY IRWIN MARTIN MARTIN HOWARD DUHUCE JOSEPH FHAZIER DOUGLAS	1969 1971 1971 1971 1972	9 - 9 9 9	6 X 6 X 6 X 6 X 6 X	150 130 102 52 130	****	22 49 12 30 3	=	20 5 F 4	 11-71 7-71 10-71 8-72	TITI	3 2 4 15 6		 1 1 1 1		· P · · ·
22223	41 42 43 44 45	442251N0723427.1 442243N0723422.1 442311N0723513.1 442306N0723517.1 442253N0723556.1	1140 1160 1340 1420 1440	NELSON JERRY FRAZIER KENNETH NYBERG JOE BLOCK JOHN HULTS WILLIAM	1973 1972 1968 1972 1970	<b>P P P P P</b>	6 X X X X X X X X X X X X X X X X X X X	63 122 220 241 82	*****	25 3 1 0 17	=	F 3 	9-73 9-72  6-72 5-70	IIIII	8 3 1 1 6		1 2 1 1	000000	
33333	46 47 48 49 50	442246N0723543.1 442244N0723538.1 442246N0723534.1 442219N0723242.1 442253N0723306.1		HULL DOUGLAS MARDEN FORD WOOTERS FRED MAXHAM DAVID J J SNACK BAR	1968 1971 1971 1964	P P P	7 X 6 X 6 0 6 -	210 62 122 118 82	*****	18 18 10 118 	=	10  60 	5-68   	TITOC	1 10 6 26		1 1 1 	D D D D D -	• • • P P
* * * *	51 1 2 3 4	442257N0723419.1 442233N0723305.1 442233N0723305.2 442304N0723304.1 4422301N0723258.1	1150 750 750 740 730	GREEN JAMES US GEOL SURVEY US GEOL SURVEY US GEOL SURVEY US GEOL SURVEY	1975 1973 1973 1974 1974	13333	02222 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	140 21 19 32 70	WTTTT	32	=	===	=	I D D D D	5		=	D DG× DG× DG× DG×	-
× × × ×	5 6 7 8	442303N0723258.1 442203N0723319.1 442219N0723242.1 442240N0723302.1	730 750 740 770	US GEOL SURVEY US GEOL SURVEY US GEOL SURVEY US GEOL SURVEY	1974 1974 1974 1974	* * * *	x x x x x x	25 87 21 62	T T T T		=		=		=	=	::	DGx DGx DGx DGx	-

#### Table 7.--Drillers' logs of selected wells and borings (Depths are given in feet below land surface.)

	De	pt	h
		Pe	
CABOT W5.			
Clay, dry	0	-	6
Sand, fine, saturated	6	-	8
CABOT W9.			10
Gravel, and dirt Hardpan (till?)	0	-	10
Rock, green	10 35	-	35 122
CABOT W21.	"		1
Sand; gravel; some brown clay			
balls	0	-	21
balls Sand, medium to coarse; gravel	21	-	26
Gravel, fine to medium; some			
coarse sand	26	-	36
Silt, gray; and very fine sand Sand, dark gray; and gravel;	36	-	37
some clay; cobbles	37	-	42.5
Boulder or bedrock	42.5	-	45.1
CALAIS W7.			
Sand, fine, brown, and very fine			
Sand, fine, brown, and very fine sand; silt; some gravel	0	-	15.1
Sand, fine, brown, and very fine			
sand; silt; trace of gravel			
some cementing of sand	15.1	-	21.2
Sand, fine, brown; very fine			
sand; silt (increasing with	21.2		37.4
depth) Sand, fine, brown; very fine	-1		57.4
sand: silt: with cemented laver	37.4	-	44.5
Sand, medium, some fine sand,			
Sand, medium, some fine sand, some coarse sand to fine gravel	44.5	-	59.5
Gravel, fine to medium, dark in			
color; some fine to medium			10
sand; broken, coarse gravel	59.5	-	65
Sand, coarse; coarse gravel; some fine to medium sand	65	-	73
Sand, fine, gray	73	-	78
Gravel, fine; and brown sand	78	-	84
Clay, gray; and fine, sharp			
gravel (till?); takes water			
rapidly at 95 feet	84	-	95.5
Silt; medium to coarse sand; gravel	95.5	-	100.7
Sand, medium to coarse, gray,			
sharp; gray wash water;			
alternately takes all water	100.7	-	110.2
Sand, fine to medium, gray to	110.0		115
green; fine, white, quartz sand	110.2	-	115
Sand, fine to medium, gray; some clay; takes most wash water	115	-	127.4
Sand, fine to medium, gray; some			
Sand, fine to medium, gray; some clay; coarse, gray sand	127.4	-	132.7
Sand, coarse, gray; and fine,			
sharp gravel; some fine, gray	122 7	-	128
sand; gray wash water Sand, fine, gray; and coarse,	132.07		150
gray sand; medium gravel;			
gray wash water	138	-	143.3
Sand, fine, gray; and medium gra- vel; sand increases with depth,			
vel; sand increases with depth,			
grains sharp and angular	143.3	-	149.4
Sand, fine to coarse, gray; some			
fine gravel; takes wash water	149.4	-	159.9
	145.4		155.5
Sand, very fine to medium, gray, brown flecks, sharp grains	159.9	-	165.3
Sand, fine to coarse, gray; fine			
sand increases with depth,			
takes water readily	165.3		
	170.6	-	1/0.3
Sand, fine to medium, gray, could not penetrate further	176.3	-	191.6
CALAIS WIG.	170.5		
Gravel, and clay	0	-	8
Sand and clay	8	-	70
Grave1	70	-	81
CALAIS W45.	0		70
Sand	0 70	-	70 76
Grave1 CALAIS X1.	10	-	10
Sand, fine to medium, yellow-			
brown; minor gravel	0	-	15.0
Sand, very fine to fine, clayey			
to silty, gray: some fine to			
medium, yellow-brown sand with gravel (probably from 0 to			
15 ft zone)	15.0	-	19.8

	De	pth	-
and a second			-
CALAIS X1 (Cont.).			
Silt, gray; minor, very fine to			
fine sand	19.8		
Clay, gray; silt; very fine sand.	24.4		
Silt and clay, gray	39.6		
Silt, clay, gray; minor very fine sand .	45.9	- 51.	
Clay, gray; silt; very fine sand.	51.2	- 61.	2
Sand, very fine, gray; silt;	61.5	- 68.	0
some clay Silt, gray; very fine sand;			-
some clay	68.0	- 78.	1
some clay Silt, gray; very fine sand; clay. Silt; very fine sand; trace of	78.1		
Silt; very fine sand; trace of			
gray clay; sand may be			
increasing with depth	83.3	- 87.	8
Silt; very fine sand; trace of			
clay; could drive casing only			
to 105 ft; washed ahead to	07.0	110	
115 ft	87.8	- 115	
CALAIS X2.	0	- 2	
Fill	0	- 3	
Organic swamp deposit, gray-	3	- 7	
black Boulders	7	- 8	
Silt brown: and gray clay.	8	- 14	
Silt, brown; gray clay: till			
Silt, brown; and gray clay Silt, brown; gray clay; till fragments; could not			
penetrate further	14	- 15	
CALAIS X3.			
Fill. Organic swamp material. Silt, brown; some very fine sand.	0	- 4	
Organic swamp material	4	- 5	
Silt, brown; some very fine sand.	5	- 15	
sill, brown; some very line			
sand; trace of pebbles	15	- 20	
Silt, gray; pebbles	20	- 25	
Silt, gray; clay Silt, gray; small flat pebbles (concretions?)	25	- 26	
Silt, gray; small flat peoples	26	- 30	
Till; could not penetrate further	30	- 30.	8
MARSHFIELD W6.	10	50.	~
Clay, blue	0	- 87	
Gravel	87	- 98	
MARSHFIELD W12.			
Clay; sand; gravel; topsoil;			
static water level at 2.0 feet			
below land surface	•	- 18.	
Clay, gray; sand; gravel	18.8		
Sand, very fine, gray; silt	24.0	- 35.	
Sand, fine, gray	35.6	- 40.	9
Sand, medium to coarse, gray to			
yellow; gravel; some silt balls; most rock fragments			
are broken granite	40.9	- 46.	1
Till, very sandy, yellow-gray;			
high percentage of granitic			
fragments in medium to coarse			
sand and gravel	46.1	- 52.	4
Till; broke off drill rod			
at 57.5 ft	52.4	- 57.	5
MARSHFIELD W35.		1.0	
Sand	0	- 40	
Clay			
Gravel; sand	50	- 63	
MARSHFIELD W36.	0	- 40	
Sand; boulders Gravel, fine	40	- 81	
MARSHFIELD W38.			
Boulders and gravel	0	- 54	
MARSHFIELD W40.			
(Sand?)	0	- 20	
Grave1	20	- 35	
MARSHFIELD W41.			
Sand, fine; boulders	0	- 45	
Gravel, fine	45	- 54	
MARSHFIELD X1.		0	
Silt, gray, wet at 6 feet	0	- 8	
Sand, medium to coarse; fine	0	- 00	
gravel; saturated	8	- 23	
Sand, very fine to medium, gray	23	- 25	
Sand, very fine to medium, gray;	25	- 30	
clay; gravel	30	- 30	
Sand, very fine, gray Clay; fine gravel	39	- 40	
Clay; fine gravel Cobbles; some sand; clay	40	- 44	
Boulders; cobbles; could not		-	
penetrate further	a	t 44	

Table	7Drillers' logs of	selected wells and borings (Co	ontinued)
		in feet below land surface.)	

		Dept	h	
		Jope		
WOODBURY W2.				
No sample; static water level at				
0.5 feet below land surface	0	-	11	
Sand, fine, gray-brown; very fine sand; some silt; traces of fine				
gravel; coarse sand; no clay	11	-	24	
Sand, medium; coarse gravel	24	-	25.5	
Sand, medium; fine gravel;				
coarse sand Sand, very fine to coarse; some	25.	5 -	30	
Sand, very fine to coarse; some	20		21	
clay; takes water readily Could not penetrate further	30	at	31 32.4	
WOODBURY W3.				
Sand, brown; fine, sharp gravel;				
sand fraction increasing				
below 16 feet	0	-	20.5	
Gravel; fine and very fine,				
gray sand; (laminated?); gravel increasing at 30-31 feet	20	5 -	31	
Sand fine dark grav: fine	20.	5 -	21	
Sand, fine, dark gray; fine pebbles; some silt; (probably				
matrix around cobbles)	31	-	34.6	
Drove casing to 33.6 ft;				
could not penetrate further;				
washed ahead		to	34.6	
WOODBURY W4.				
Sand, medium to coarse, brown;				
static water level at 2.5 feet.	0	-	25	
some silt; gravel; cobbles; static water level at 2.5 feet. Sand, coarse; broken cobbles	25	-	30	
Cobbles, broken, black and				
white; some sharp gravel	30	-	35	
Till; could not penetrate further	35	-	39.8	
WOODBURY W6.	0		51	
Gravel, fine; fine sand Coarse gravel	v	at	51	
WOODBURY W11.		ac	21	
Sand	0	-	80	
Shale (?); grave1	80		110	
Sand	110	-		
Grave1	140	-	157	
WOODBURY W12.	0		20	
Sand	0 30	- 2	30 48	
Gravel, fine	30		40	
Sand, brown; coarse gravel	0	-	16	
Drove casing to 20.5 feet; could				
not penetrate further		at	20.5	
WOODBURY X2.			0 -	
Fill, roadway; till Could not penetrate further	0	-	8.5	
WOODBURY X3.		at	0.5	
Fill, roadway; boulders; till	0	-	21.4	
Could not penetrate further		at	21.4	
WORCESTER WI.				
Sand, brown; gravel	0	-	18	
Clay, gray. Clay, gray; silt; artesian flow at 37.6 feet	10	-	32.4	
at 37.6 feet	32.	4 -	37.6	
Silt, gray	37.	.6 -	53	
Silt, gray; sand	53	-	54	
Gravel, fine; gray silt	54	-	63.4	
Could not penetrate further			63.4	
(bedrock or boulder) WORCESTER W2.		at	03.4	
Sand; grave1	0	-	12	
Sand, fine (muscovite flakes),	-			
grades to gray clay	12	-	20	
Clay, gray	20	-	34	
Clay, gray. Sand, fine, gray; some fine gravel at top; static water				
level at 3.5 feet	34	-	41	
Sand, fine gray; gray clay layers	41	-	52	
Clay, gray	52	-	67.5	
Sand, fine	67	.5 -	68	
Clay, gray	68	-	75	
Sand, fine and medium, gray; some				
coarse sand; static water level	75		80	
at 14 feet below land surface Sand, fine and medium, gray; and	75	-	80	
clay layers	80	-	85	
Sand, fine, clayey, gray; some				
coarse sand	85	-	90	
Sand, fine, clayey, gray; trace				
of gravel Sand, fine, clayey, gray; some	90	-	95	
coarse sand; very fine gravel	95	-	112	
source sand, very time graver	,,			
				-

		Dept	h
		- op c	
WORCESTER W2 (Cont.).			
Sand, fine, clayey, gray; coarse			
sand increasing; some fine gravel at 120 ft; static water			
level 1 ft above land surface	112	-	120
Sand, fine to coarse, sharp; some			
gravel; clay; static water level	100		120
21.4 ft below land surface Sand, coarse, sharp, gray; silt;	120	-	130
clay	130	-	132
Sand, coarse, gray	132	-	135
Sand, coarse, gray Weathered rock fragments; sharp, gray to green (bedrock or boulder)	1.00		1.04
to green (bedrock or boulder)	135	-	136
WORCESTER W12. Clay, blue	0	-	20
Silt; fine sand; clay	20	-	55
Gravel, fine	55	-	61
WORCESTER W15.			
Gravel; fill	0	-	5 95
Clay	95	-	101.5
Gravel			10115
Alluvium	0	-	17.4
Clay, gray; sand; rock fragments. Could not penetrate further	17.	4 -	20.7
Could not penetrate further			20 7
(bedrock or boulder) WORCESTER X2.		at	20.7
Alluvium; sand; gravel	0	-	14
Clay, gray	14	-	16.5
No sample	16.	5 -	19.4
Could not penetrate further			19.4
(bedrock or boulder) WORCESTER X3.		at	19.4
Sand, brown; silt; pebbles;			
static water level about 10 ft.	0	-	10
Silt, gray; some clay	10	-	15
Clay, gray	15	-	30
Clay, gray; rock fragments Could not penetrate further	30	-	31.5
(bedrock or boulder)		at	31.5
WORCESTER X4.			
Sand, brown; gravel	0	-	9
Silt, gray	9	-	15
Clay, gray. Clay, gray; some sharp pebbles	15 20	-	20 25
Silt, gray; clay	25	-	39
Silt, gray; clay; some pebbles	39	-	45
Silt, gray; clay	45	-	50
Silt, gray; clay; some pebbles	50	-	55 60
Silt, gray; trace of clay Silt, gray	55 60	-	69.7
Could not penetrate further	00		03.1
(bedrock or boulder)		at	69.7
WORCESTER X5.			
Silt, brown; coarse sand; saturated.	07	-	7
Gravel, coarse, brown Clay, gray; stopped at 25 ft	9	-	25
WORCESTER X6.	-		
Gravel, brown	0	-	8
Sand, fine, brown Clay, gray; dry	8	-	15
Clay, gray; dry	15	-	21
Clay, gray; silt; very fine sand; layered	21	-	25
Clay gray	25	-	50
Silt, gray; some clay Sand, very fine, gray; some	50	-	75
Sand, very fine, gray; some	70		00
silt; fine sand Sand, very fine, gray; trace	75	-	80
of gravel	80	-	85
Till	85	-	86.6
Could not penetrate further			
(bedrock or boulder) WORCESTER X7.		at	86.6
Clay, gray; dry	0	-	3
Clay, gray; wet	3	-	21
Clay, gray; wet Could not penetrate further	-		
(bedrock or boulder)		at	21
WORCESTER X8.		1	14
Sand; gravel; wet at 9 feet Silt, brown	0	-	14 19
Silt, gray		-	20
Silt, gray Clay, gray and brown; laminated	20	-	25
Clay, gray; silt	25	-	45
Clay, gray; trace of pebbles	45 55	-	55 61.5
Silt, gray-green Could not penetrate further	22	1	01.5
(bedrock or boulder)		at	61.5

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