

INTRODUCTION Location and General Features

A gravity study has been made of the Twentynine Palms area covering about 375 mi² (square miles) of the Twentynine Palms Marine Corps Base and vicinity, Calif. The main facility of the Twentynine Palms Marine Corps Base is about 1,000 mi² (square miles) east of Los Angeles and 5 mi north of Twentynine Palms. Towns in or adjacent to the study area are Twentynine Palms, Joshua Tree (off map), and Landers. The study area covers about 210 mi² in the south-west corner of the base. The area adjacent to the base is rural to undeveloped. The base is completely undeveloped except for the main facility and a water-well field at Surprise Spring, and is used principally for military exercises and training. The study area is a large sedimentary filled valley surrounded by numerous hills and mountains of igneous and metamorphic rocks (Precambrian and Mesozoic age), and is crossed by several large and many small faults. The geology of the area has been mapped by Dibblee (1967a, b, c, and 1968) and is shown in simplified form on the gravity map. Four of the large faults on the base (Emerson-Copper Mountain, Hidalgo-Surprise Spring, Calico, and Mesquite) act as ground-water barriers. Water-level measurements (Bader and Moyle, 1960) show that the Surprise Spring fault is at land surface west of the fault, and 300 ft below land surface east of the fault. Several of these ground-water barriers have been used as boundaries for ground-water basins by various investigators such as Schaefer (1978).

Purpose and Scope of Work

The entire Marine Corps Base water supply is dependent on ground water. Water levels in wells near Surprise Spring have declined a maximum of 100 ft between 1950 and 1982. Because of the water-level declines and possible expansion of base facilities, the U.S. Marine Corps, Department of the Navy, asked the Geological Survey to make a multiphase study of the feasibility for basin recharge with imported water. The purpose of the gravity survey, the first phase of the multiphased study, is to determine the thickness of the sedimentary deposits that contain the bulk of the water within the study area. A map showing lines of equal sedimentary thickness was constructed from 495 gravity measurements obtained in the study area. This work included detailed measurements along profiles across areas where suspected ground-water recharge enters the study area and adjacent areas. It also included a profile across the ground-water outflow section from the study area. In all, 13 profiles were made to construct the thickness map, and all these profiles have been modeled using two-dimensional computer programs developed by Talwani (1968).

Previous Work and Acknowledgments

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GRAVITY SURVEY

Gravity surveys measure differences in the attraction of gravity at the Earth's surface caused by differences in density of the material beneath the surface. The differences may be used to indicate geologic structure and general shape of the basement complex beneath the sedimentary deposits. Because of some density variations between rock types, the complete Bouguer gravity anomaly map should not be interpreted as indicating depth to basement complex directly. In general, a gravity low corresponds to a thick sedimentary section and a gravity high corresponds to a thin sedimentary section.

Field Work

Throughout the study area during 1981-82, gravity measurements were made about 1 mi apart; however, some measurements along profiles are closer to give greater detail, and some are farther apart where access into the area is difficult. All measurements were made with Warden gravity meters (No. 782 and No. 1083) having meter constants of 0.45779 and 0.20296 mgal (milligals) per scale division, respectively. A few measurements on the northern edge of the study area, made by Biehler and others (1984), were used for this report. All gravity measurements made for this study were tied to a local base station¹ at the northeast corner of Valle Vista Road and Utah Trail, which is 1 mi east and 4 mi north of the center of Twentynine Palms near the main entrance to the Marine Corps Base. This local base station was established relative to the California Base Station Network by tying to the California Base Station Network at Amboy, Calif. (Chapman, 1966, Station 292), and is thus on the Woodard and Rose (1963) gravity datum. The error of closure between the Amboy station and the local base station at Twentynine Palms was 0.1 scale division or 0.02 mgal.

¹The use of the brand name is for identification purposes only and does not constitute endorsement by the U.S. Geological Survey.

²Base station is located at a bench mark about 25 ft northeast of the intersection of Valley Vista and Utah Trail in T.2 N., R.3 E., sec. 3, San Bernardino base line and meridian.

Reduction of Data

Data were corrected manually for meter drift and terrain effects to 2.29 km (kilometers) using tables developed by Hayford and Bowie (1912). The data were then entered into a computer and corrected for latitude, altitude (combined free air and Bouguer effects), Earth curvature, and terrain effects from 2.29 to 90 km. A density of 2.67 g/cm³ (grams per cubic centimeter) was used for the Bouguer reduction. Terrain corrections for all regions from 2.29 km to 90 km from each station were made using Plouff's (1965) computer program and previously digitized topography from Robbins and others (1973). The resulting computer readout lists the complete Bouguer gravity anomaly for each station which was then used to construct the contours. To eliminate negative values, 1,000 mgal were added to each Bouguer anomaly value. Thus, for example, the true Bouguer anomaly corresponding to the 910 contour shown on the map is -90 mgal.

Bouguer Gravity Anomaly Map

The Bouguer gravity anomaly map was contoured from the plotted data points. The gravity values (Bouguer anomaly +1000) range from a high of 925 mgal in the mountains at the northwest corner of the study area to a low of 880 mgal in the northeast part of the study area.

Included on the gravity anomaly map are the geologic contacts between the basement complex and sedimentary deposits and the locations of faults (after Dibblee, 1967a, b, c, and 1968), and a diagram showing rock densities. The bars on the diagram show the upper and lower limits of rock densities measured by Mabey (1960) in his gravity survey of the western Mojave Desert. The circles show the measured densities of rocks from the study area. Data from 17 rock samples indicate that the basement complex in the Twentynine Palms area has an average density of 2.49 g/cm³. The sedimentary deposits have an average density of 2.15 g/cm³, and a density contrast of 0.34 g/cm³. Based on these limited number of samples, the computed depth to basement complex using 0.34 g/cm³ agrees favorably with depth to basement complex shown on well logs between 400 and 2,100 ft deep.

Interpretation of the Gravity Data

To interpret the data, the regional gradient must be removed from the Bouguer gravity map. A computer program written by Jachens and Roberts (1981) was run on the U.S. Geological Survey computer in Menlo Park, Calif. to remove the regional gradient due to isostatic compensation of topography. This program did not remove the gravity variations caused by density variations within the basement complex. Jachens (U.S. Geological Survey, oral communication, 1982) suggested that the regional gradient variations from this source be removed as a single correction over each short profile between basement complex outcrops. The regional gradient along each profile was removed before two-dimensional modeling was done. Thirteen gravity profiles were modeled between basement complex outcrops in the study area. Each profile was made assuming a density contrast of 0.34 g/cm³. After modeling was completed, the calculated depths to basement complex using 0.34 g/cm³ agrees favorably with depth to basement complex shown on well logs located in the same area. Although many wells have been drilled in the study area, only three wells reach basement complex near where profiles were made. Well 2N/7E-26B1 (map No. 1) has a computed depth of 400 ft and a well log depth of 405 ft. Well 2N/6E-321 (map No. 2) has a computed depth of less than 500 ft and a well log depth of 149 ft. One oil test hole, 2N/8E-280 (map No. 27), was drilled to a depth of 2,106 ft. Data in the California Division of Oil and Gas Report (1964, p. 55) show this well encountered "granite basalts" at the bottom of the hole. This may indicate decomposed granite in the lower part of the well. From this description the exact depth to basement complex could not be determined; however, the gravity data indicate the depth to basement complex is between 1,500 and 2,000 feet. All other computed depths are deeper than the total depth of the wells in the same areas. Thus, there is some rough confirmation of depths to basement complex computed from gravity data from a few shallow wells, but there is no check on computed depths greater than 2,000 ft. The supplemental well data show the thickness of material penetrated at each well. After profiling was completed, lines of equal thickness of sedimentary deposits were added to the Bouguer gravity map by contouring the calculated depths along the profiles.

Error Analysis

Zbur (1963) shows that sedimentary deposits in desert basins become more dense with depth, and the density contrast between basement complex rocks and sedimentary deposits generally range between 0.5 and 0.3 g/cm³. Bostain and others (1976) state that in general, "A density contrast of 0.5 g/cm³ might better fit basins less than 2,000 ft deep. A density contrast of 0.4 g/cm³ will best fit basins between 1 and 2 km deep, and a density contrast of 0.3 g/cm³ best represents deeper basins." For this study, 0.34 g/cm³ was used for all profiles, producing very good results in areas where depth to basement complex (based on well logs) is 0.6 km (2,100 ft) or less. To compute possible errors introduced because of incorrect average bulk density contrast values, densities of 0.3 and 0.4 g/cm³ were used on the profile that crosses the valley at the north end of Deadman Lake at the largest gravity anomaly in the study area. Using these values, the model shows that if 0.3 g/cm³ is used, the maximum thickness of sedimentary fill would be about 15,500 ft, and if 0.4 g/cm³ is used, the minimum thickness of sedimentary fill would be about 8,300 ft. Using 0.34 g/cm³, the thickness of sedimentary deposits is about 10,300 ft. Since the results using 0.34 g/cm³ in areas where depth to basement complex is 0.6 km (2,100 ft) deep or less correlate so well with well log depths, it would seem reasonable that the deep part of the basin has a density contrast of 0.34 g/cm³ or less. This indicates the sedimentary deposits in the basin are at least 10,300 ft thick. Some two-layer profiles were also modeled along this same profile in an error-analysis procedure. These profiles could not be used because of the lack of well log data to verify the assumptions used in construction of the profiles; however, they did not drastically differ from the single layer model.

CHEMICAL ANALYSES OF WATER

The gravity data show that areas of good and poor quality water are separated by a buried bedrock ridge southeast of Emerson Lake and northeast of the Emerson fault. Gravity station 297 shows this ridge to be about 250 ft below land surface. Chemical analysis of water from three wells near Ames Lake in T.4 N., R.6 E., sec. 27, show large differences in amounts of various constituents dissolved in the water (Riley and Bader, 1961, p. 72). This is probably due to well depth, which ranges from 63 to 182 ft. The dissolved solids in these wells do not meet the recommended criteria for drinking water. Other wells near

Ames Lake and Emerson Lake also contain poor quality water. Between Ames Lake and Surprise Spring the data show that all wells contain water of good quality, ranging in dissolved solids from about 100 to 200 mg/L.

SUMMARY AND CONCLUSIONS

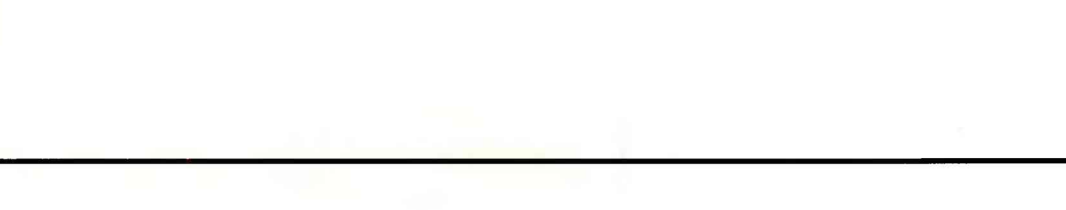
The sedimentary deposits that overlie the basement complex in the study area range in thickness from less than 1 ft at the edge of the valley at the basement complex outcrop, to a calculated depth of 10,300 ft near the center of the gravity low in the northeast part of the study area. This estimated depth may be as much as 25 percent greater in areas where the depth to basement complex is greater than 2,000 ft if the density contrast is less than the 0.34 g/cm³ used in the model. The map shows four areas of thick alluvial deposits beneath the Marine Corps Base. The area west of the Mesquite fault near the Marine Corps Training Center headquarters in T.2 N., R.3 E., has a maximum calculated thickness of sedimentary deposits of 4,700 ft. In the south-western corner of the base, the maximum calculated thickness of sedimentary deposits is about 2,300 ft. In the area near the western edge of the base southeast of Emerson Lake, two thick sedimentary sections exist on each side of the Emerson fault. The section on the northeast side of the Emerson fault near Ames Lake is about 1,000 ft thick, and on the southwest side is about 2,700 ft thick. Some faults act as barriers to ground-water movement in sedimentary deposits (Bader, 1963). The surface expression of these faults has been mapped by Dibblee (1967a, b, c, and 1968). The barrier effect of these faults at the water table may be shown at a different location than the surface expression if the dip of the fault is not vertical. Where known, faults and ground-water barriers are both shown on this map. The map also shows areas where the basement complex has an altitude about the same as that of the water table in the sedimentary deposits. These areas have poor potential for developing production wells for water supply.

The most important use of this map will be in later phases of the overall study, when the map will be used to determine the amount of ground water in storage in the individual basins.

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Supplemental Well Data table with columns: Map No., State well No., Thickness of sedimentary deposits in well in feet, Depth to basement complex in feet, Depth of well in feet.



CONTOUR INTERVAL 40 FEET NATIONAL GEODETIC VERTICAL DATUM OF 1929

BOUGUER GRAVITY ANOMALY MAP OF THE TWENTYNINE PALMS MARINE CORPS BASE AND VICINITY, CALIFORNIA SHOWING THICKNESS OF SEDIMENTARY DEPOSITS, DEEP WELLS, AND GEOLOGY

By W. R. MOYLE, JR. 1984

Black and white copies of this map may be purchased from: For additional information write to: Open-File Section, Western Distribution Branch, U.S. Geological Survey, Box 2425, Federal Center, Denver, Colorado 80225 (Telephone: 303/234-6888) District Chief, U.S. Geological Survey, Federal Building, Room W-2225, 2601 Cottage Way, Sacramento, CA 95833