

PROJECTED EFFECTS OF GROUND-WATER WITHDRAWALS IN THE ARKANSAS RIVER
VALLEY, 1980-99, HAMILTON AND KEARNY COUNTIES, SOUTHWESTERN KANSAS

By L. E. Dunlap, R. J. Lindgren, and J. E. Carr

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CONVERSION TABLE

Inch-pound units of measurement in this report may be converted to International System (SI) of Units using the following conversion factors:

<u>To convert from</u> <u>inch-pound units</u>	<u>To SI</u> <u>unit</u>	<u>Multiply by</u>
inch	millimeter	25.4
foot	meter	0.3048
mile	kilometer	1.609
acre	square meter	4,047
square mile	square kilometer	2.590
acre-foot	cubic meter	1,233
cubic foot per second (ft ³ /s)	cubic meter per second	0.02832
foot per day (ft/d)	meter per day	0.3048
acre-foot per year (acre-ft/yr)	cubic meter per year	1,233
inch per year	millimeter per year	25.4

DEFINITION OF TERMS

Aquifer - A formation, group of formations, or part of a formation that contains sufficient saturated permeable material to yield significant quantities of water to wells or springs.

Consumptive use - Volume of water that is used by vegetative growth in transpiration and building of plant tissue and that is evaporated from adjacent soil or intercepted precipitation on plant foliage.

Deep percolation - Volume of water from precipitation and irrigation that infiltrates the soil and moves by the force of gravity to the water table.

Evapotranspiration - Volume of water that is lost to the atmosphere by transpiration from vegetative growth and by evaporation from the soil or from the aquifer in shallow water-table areas.

Head, static - The height above a standard datum of the surface of a column of water than can be supported by the static pressure at a given point.

Hydraulic conductivity - Volume of water at the existing kinematic viscosity that will move through a porous medium in unit time under a unit hydraulic gradient through a unit area measured at right angles to the direction of flow.

Hydraulic gradient - The change in static head per unit of distance in a given direction.

Moisture-holding capacity - Amount of moisture that the soil can hold in a form available to plants. It is the amount of moisture held between field capacity and the permanent wilting point.

Specific yield - Ratio of the volume of water that the saturated material will yield by gravity drainage to the volume of the material.

Steady state - Equilibrium conditions when water levels and the volume of water in storage do not change with time.

Storage coefficient - Volume of water an aquifer releases from or takes into storage per unit surface area of the aquifer per unit change in head.

Streambed leakance - Ratio of the vertical hydraulic conductivity of the streambed to the thickness of the streambed material.

Transmissivity - Rate at which water of the prevailing kinematic viscosity is transmitted through a unit width of an aquifer under a unit hydraulic gradient.

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ABSTRACT

A study was made, in cooperation with the Division of Water Resources, Kansas State Board of Agriculture, to determine the effects that additional ground-water development would have on streamflow and water levels in an area along the Arkansas River in Hamilton and Kearny Counties, southwestern Kansas. A computer model was used to simulate the changes in streamflow and water levels from 1980 through 1999. Six pumpage options were tested using variations in pumpage rate and number of wells pumping in the model area.

The projected effects of ground-water withdrawals from 1980-99 indicate that net annual streamflow losses would be reduced only 1 percent if annual withdrawals were reduced by 24 percent from the continued 1979 pumpage rate of wells with water rights, as computed using energy-consumption techniques, to the amount appropriated by water rights for these wells. The higher the pumpage rate results in very little change in stored ground water and net streamflow losses because, after satisfying initial soil-moisture requirements, a greater percentage of the additional applied water percolates back to the aquifer.

If pumpage were increased along with an increase in the number of irrigated acres over the 1979 rate, additional water would be removed from storage, and net annual streamflow losses would increase. A 19-percent increase in pumpage over the 1979 rate, with the additional acres irrigated, would cause the net annual streamflow loss to increase 5 to 9 percent.

INTRODUCTION

In January 1977, the Chief Engineer of the Division of Water Resources, Kansas State Board of Agriculture, declared a moratorium on the approval of applications for permits to appropriate water from an area of 500 square miles along the Arkansas River in Hamilton and Kearny Counties, southwestern Kansas (fig. 1). The moratorium was prompted by a growing concern over decreasing streamflow and declining water levels and by the need for a better understanding of ground- and surface-water interaction in the area.

When the moratorium was declared, hydrologic information was insufficient to allow adequate scientific evaluation of the interaction of ground and surface water and the extent to which diversion from either source might impair water use under existing water rights in the area. The effect of year-to-year decreases in streamflow at the State line during

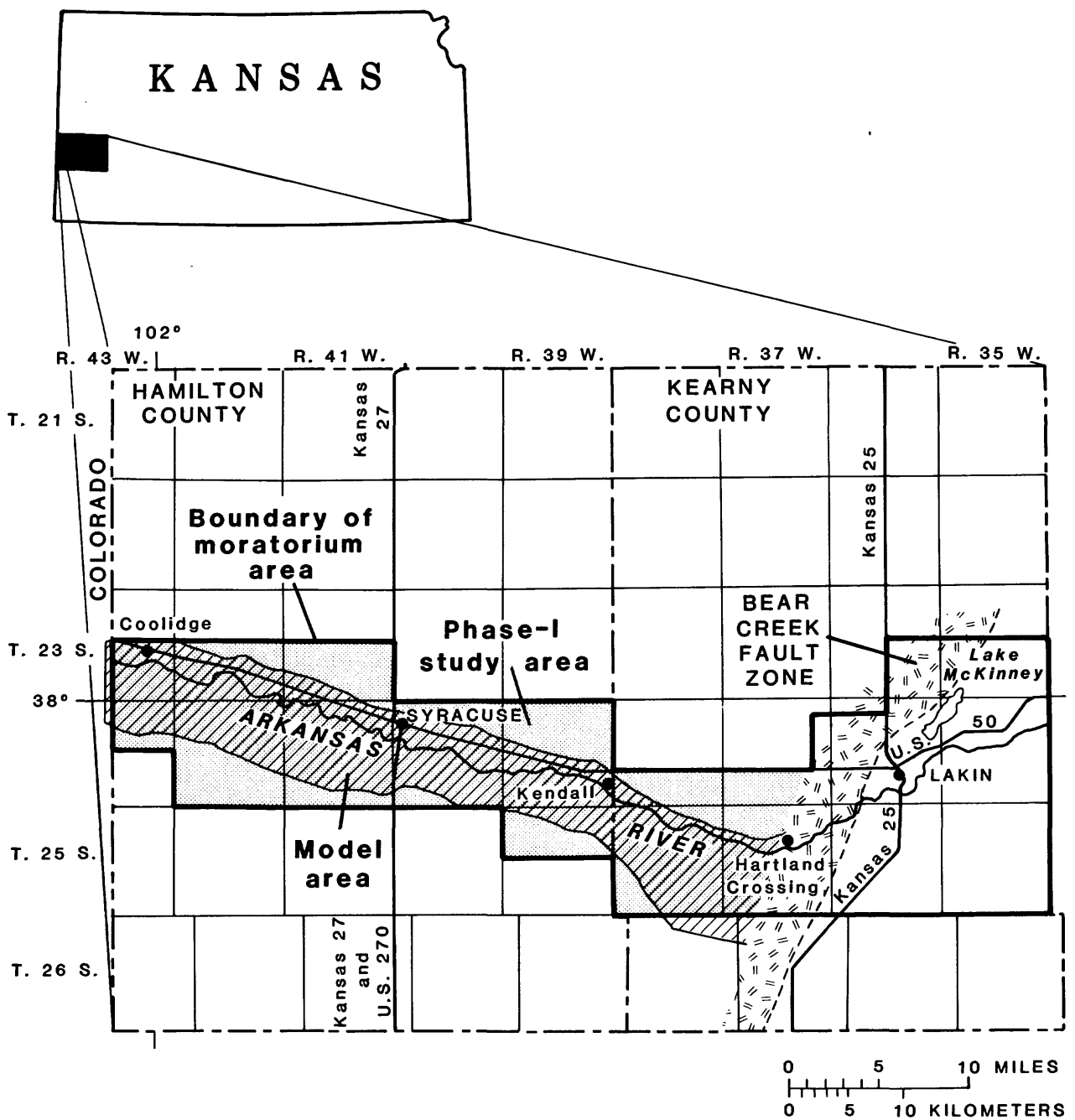


Figure 1.--Location of moratorium, study, and model areas.

the 1970's also needed evaluation. The Division of Water Resources entered into a 5-year cooperative investigation of the moratorium area with the U.S. Geological Survey in October 1977 to provide better hydrologic information for use in managing the water resources of the area.

The study of the moratorium area was divided into two parts in order to separate the geological differences due to the Bear Creek Fault zone--phase I (1977-81) and phase II (1979-82). In 1981, the primary investigation of the stream-aquifer hydrology along the Arkansas River valley west of the Bear Creek Fault zone (phase I) was completed (Barker and others, 1983). As part of the phase-I study, long-term model projections were used to assess the effects that additional pumpage might have on future streamflow and water levels (fig. 2).

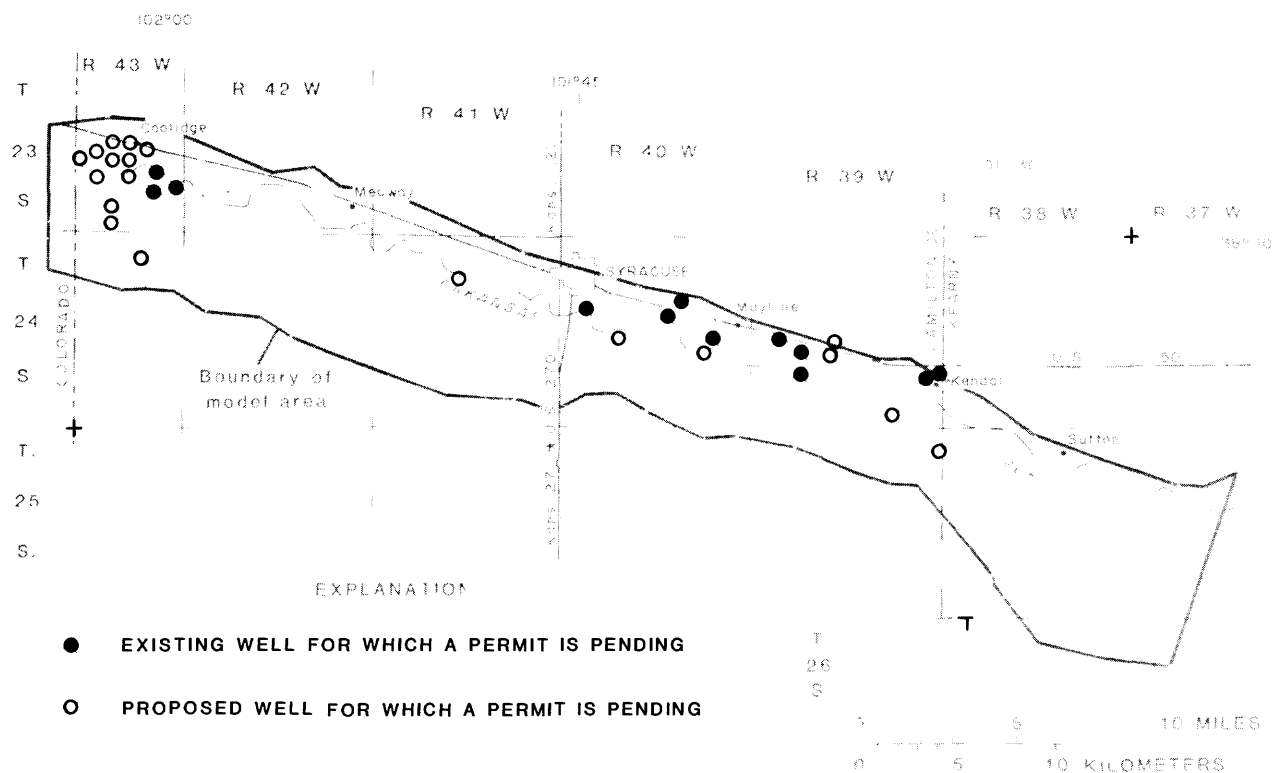


Figure 2.--Location of additional pumpage sites used in model projections.

Purpose and Scope

The purpose of this report is to evaluate the effects that additional ground-water development might have on streamflow loss in the Arkansas River and on water-level changes in the aquifer from 1980 through 1999. Sixteen long-term projections of possible future hydrologic conditions were made using a calibrated finite-element model. The report will describe briefly: the hydrologic properties used as input to the model, the 16 projection options used, and the results of those projections. For complete information on the environmental setting of the area and the background, input, and calibration of the computer-model analysis, the reader is referred to a previous report of the study area (Barker and others, 1983). Documentation for the Tracy finite-element model used to make the long-term projections is presented at the end of this report.

Method of Investigation

Digital-Computer Model

The digital model of the Arkansas River valley stream-aquifer system uses a computer program written by J. V. Tracy (U.S. Geological Survey, written commun., 1980). The Tracy finite-element model is based on the Galerkin finite-element method (Desai and Abel, 1972). Additional reference to this method can be found in Pinder and Gray (1977). A "direct" solution method is used to solve the nonlinear, partial-differential equation that describes nonsteady, two-dimensional ground-water flow.

The simulation of future conditions requires that time be broken into a series of finite intervals called time steps. Because the solution changes with time, the size of the time steps affects the computational-work effort needed to approximate aquifer performance during a selected period and the accuracy of the approximation. A progression of 5-day time steps was used for all model analyses.

Development of the digital-computer model as a predictive tool is based on the premise that, if historic hydrologic phenomena can be satisfactorily approximated by the model, then so can future conditions. The historic cause-and-effect relationship between stresses in the real flow system and the system's response to those stresses were simulated with acceptable accuracy in the previous study by Barker and others (1983). The investigation reported here assumed that the cause-and-effect relationship did not change significantly in the real system during the simulation period, 1980 to 1999.

However, the hydrologic system in the study area is very dynamic. Large changes in streamflow at the Colorado-Kansas State line and in ground-water pumpage during the 1970's greatly influenced streamflow and water levels in the study area (Barker and others, 1983). Although it may be possible to estimate future pumpage based on management control, it is highly speculative to forecast changes in the streamflow at the State line for the 1980's and 1990's. The method of investigation chosen was to input various streamflow and pumpage patterns, allowing interested authorities to make management decisions based upon the different results.

Acknowledgments

The staff of the Division of Water Resources, Kansas State Board of Agriculture, especially David Pope, Chief Engineer, and Howard Corrigan, Water Commissioner, are thanked for the information they provided on water rights, surface-water diversion, well discharge, and water levels for the study.

SIMULATED HYDROLOGIC PROPERTIES

Boundary conditions on the east and west sides of the model area, precipitation, streamflow conditions, and pumpage conditions were the only simulated hydrologic properties whose values for the 1980-99 projections varied from the values used in the previous study by Barker and others (1983).

Boundary Conditions

Boundary conditions were specified in order to terminate lateral ground-water flow at the northern and southern boundaries of the modeled system and to maintain the December 31, 1979, water level at the western and eastern boundaries. The boundaries on the north and south depicted the effects of the thinning of the alluvial aquifer system to termination against the relatively impervious bedrock. The boundaries on the east and west maintained constant water levels at nodes representing lateral inflow across the Colorado-Kansas State line and lateral outflow near Hartland Crossing. Specified water levels at these locations were maintained at the December 31, 1979, altitude throughout the 1980-99 projections because of the unpredictability of water-level trends in the distant future. In the previous study by Barker and others (1983), specified hydraulic-head nodes, which were updated at the beginning of each time step, were used on the east and west boundaries of the modeled area. The simulated water-level surface on December 31, 1979, was used as the starting water-level surface for the 1980-99 projections.

Streamflow Conditions

The model used 87 riverbed nodes to control the simulation of stream-aquifer interaction and to orient the routing of river discharge from the Colorado-Kansas State line to the downstream limits of the modeled area. Streambed altitude, lengths, and widths were specified for each riverbed node as part of the model input. The model simulated rates of streamflow for all riverbed nodes in an iterative, upstream-to-downstream fashion. Starting with input streamflow at the uppermost stream node and working downstream, the flow was calculated for each reach on the basis of incoming flow and the gain from, or loss to, the aquifer throughout the length of the reach.

Variations in 1971-80 streamflow at the Colorado-Kansas State line were used in running the 1980-99 future projections. Monthly discharge of the Arkansas River at the Colorado-Kansas State line for a U.S. Geological Survey streamflow-gaging station (near Coolidge) during 1970-80 is shown in figure 3. The average annual discharge during 1951-69 is also shown. Streamflow during the early 1970's was much higher than during the late 1970's. The streamflow during 1980 was exceptionally high.

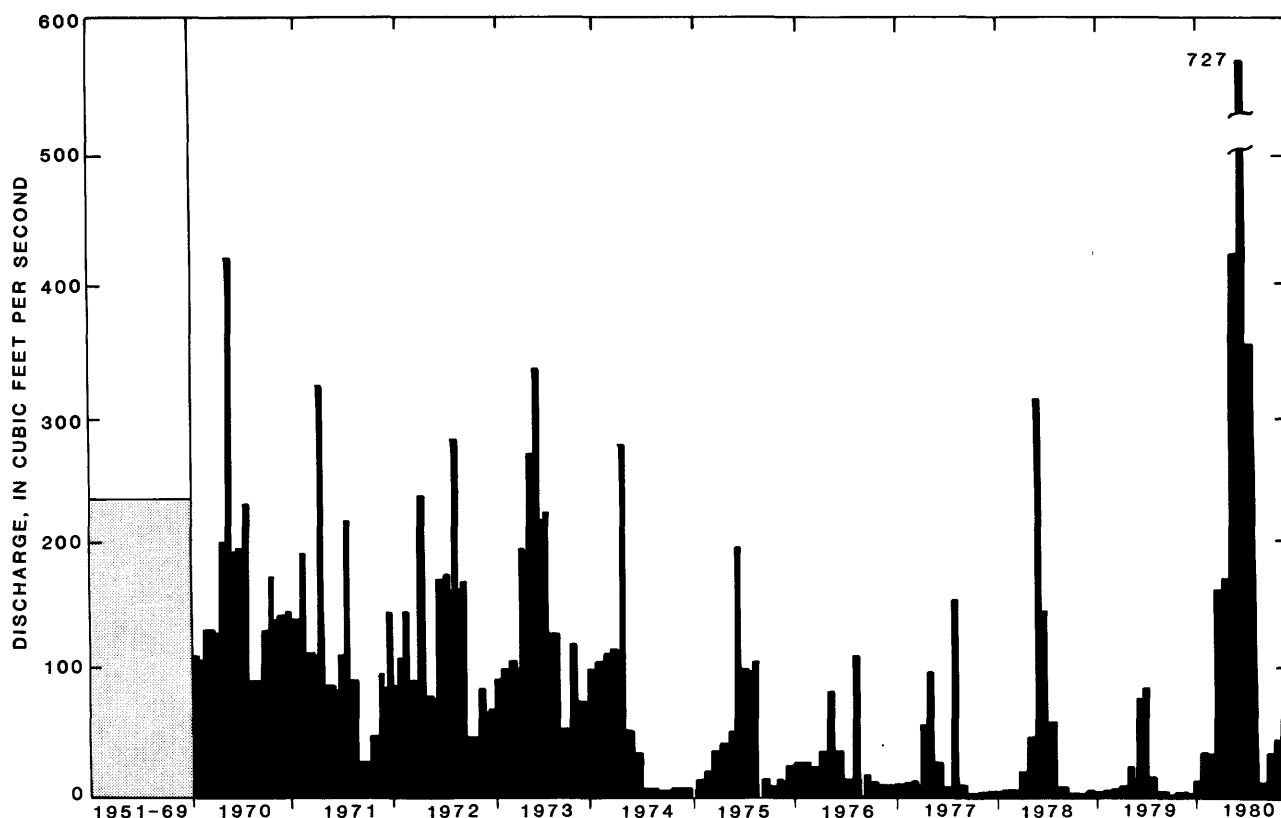


Figure 3.--Discharge of Arkansas River near Coolidge, 1951-69 (average) and 1970-80 (monthly).

Four variations in streamflow conditions were simulated:

- OPTION A - Cycle the 1971-80 gaged daily flow of the Arkansas River near Coolidge for the consecutive time periods 1980 through 1989 and 1990 through 1999. In other words, the gaged daily flow near Coolidge on January 1, 1971, was used for the streamflow on January 1, 1980, and January 1, 1990. The dates correspond until gaged daily flow on December 31, 1980, was used for the streamflow on December 31, 1989, and December 31, 1999.
- OPTION B - Cycle the 1971-80 average monthly flow of the Arkansas River near Coolidge for the time period 1980-99. In other words, averages of the individual months during the 1971-80 time period were used as daily flow from 1980 through 1999.
- OPTION C - Cycle the 1976-80 gaged daily flow of the Arkansas River near Coolidge for the consecutive time periods 1980-84, 1985-89, 1990-94, and 1995-1999. This is similar to option A, except that the gaged daily flow near Coolidge on January 1, 1976, was used for the streamflow on January 1, 1980, January 1, 1985, January 1, 1990, and January 1, 1995.
- OPTION D - Cycle the 1979 gaged daily flow of the Arkansas River near Coolidge throughout the 1980-99 projection.

Aquifer Parameters

Values for hydraulic conductivity, a measure of the aquifer's ability to transmit water, and saturated thickness, which is dependent on water levels, were used by the model to compute the required transmissivity distribution during the simulation. An area-constant hydraulic conductivity of 800 ft/d was used in the calibrated model for the Arkansas River alluvium. This was the same value used in the investigation by Barker and others (1983, p. 25).

The magnitude of water-level change that occurs in a water-table aquifer in response to recharge or discharge of ground water depends on the specific yield. The Arkansas River model used a simple distribution of specific yield that ranged in value from 0.14 to 0.20. This was the same distribution used by Barker and others (1983, p. 26).

Discharge

Pumpage and Pumpage Options 1-6

The largest vertical discharge from the aquifer was pumpage. The pumpage history of the modeled area was estimated by making an inventory of public-supply and irrigation wells in the area and by computing monthly pumpage rates for those wells from energy-consumption and, in a few instances, water-right and billing records. Area-wide averages were determined from monthly pumpage rates, where available, and applied to annual totals for other wells to provide the monthly patterns of pumpage throughout the area (Barker and others, 1983). Variations in these pumpage estimates and the appropriated amount on the landowner's water right were used in the 1980-99 projections.

Six hypothetical pumpage conditions were used in the 1980-99 projections. They were:

- OPTION 1 - Includes pumpage from all wells in the study area (147 wells as of 1981) that had water rights from the Division of Water Resources. Pumpage rates were determined from 1979 energy-consumption, water-right, or billing records. The 1979 calculated ground-water withdrawals are greater than the appropriated rights when this technique is used to determine pumpage.
- OPTION 2 - Includes pumpage from all wells in the study area that had water rights as of 1981. Pumpage rates were determined from the amount of water appropriated on the water rights. This option includes 152 wells. Five additional wells were added, in addition to those in option 1, that did not pump during 1978 or 1979.

- OPTION 3 - Includes pumpage from all wells that had water rights, plus pumpage from 12 existing wells for which permits were pending (159 wells as of 1981). Pumpage rates for all these wells were determined from 1979 energy-consumption, water-right, or billing records.
- OPTION 4 - Includes pumpage from all wells that had water rights, pumpage from existing wells for which permits were pending, plus pumpage from an additional 19 proposed wells for which permits were pending (178 wells as of 1981). Pumpage rates for wells with water rights and existing wells for which permits were pending were determined from 1979 energy-consumption, water-right, or billing records. Pumpage rates for proposed wells for which permits were pending were based on the amounts of water requested by well applications.
- OPTION 5 - Includes the same wells as in OPTION 4 but use pumpage rates determined from the amount of water appropriated on the water right or requested by well application. This option includes 183 wells. Five additional wells were added, in comparison to to option 4, that did not pump during 1978 or 1979.
- OPTION 6 - Includes the same wells as in OPTION 3 but increase the 1979 pumpage rate of each well by 50 percent (159 wells).

Monthly patterns of the six hypothetical annual pumpage conditions were cycled 20 times over the 20-year projection period.

Stream-Aquifer Leakage

The exchange of water between the Arkansas River and the alluvium occurs through the streambed. Stream-aquifer leakage can be either a source of recharge or discharge to the aquifer. A streambed-leakance value of 1.34 d^{-1} was used for all riverbed nodes in the model (Barker and others 1983, p. 30).

Ground-Water Evapotranspiration

Significant discharge can occur from the Arkansas River alluvial aquifer as ground-water evapotranspiration when the water table is above the root zone or within reach of roots through capillary attraction. Ground-water evapotranspiration can completely satisfy the consumptive-use demand by plants if the water table is at or above land surface. That part of the deficit that can be met when the water table is below land surface declines by about 8 percent per each additional foot below land surface, reaching zero at depths of 12 feet and greater. Ground-water evapotranspiration was calculated by the model in the same manner as in the previous investigation (Barker and others, 1983, p. 34).

Recharge

Deep Percolation

Recharge to the ground-water system occurs as water infiltrates from the land surface through the soil zone to the aquifer. The sources of water that may infiltrate from the land surface are precipitation and irrigation water (including both well pumpage and surface-water diversion). The amount of deep percolation depends on the amount of precipitation and irrigation water applied to the land surface, the rate of consumptive-use demand by plants, and the moisture-holding capacity of the soil zone. The model used a composite area-weighted average of crop and land-use categories to determine the consumptive-use demands (Barker and others, 1983).

Precipitation

When precipitation exceeds storage capacity of the soil, it recharges the aquifer as part of deep percolation. Precipitation in the moratorium area ranges from about 14.5 in/yr near the Colorado-Kansas State line to about 17.5 in/yr at Lakin. The 30-year (1941-70) normal precipitation at Syracuse, near the center of the modeled area, is 16.86 in/yr. The monthly normal precipitation at Syracuse was used in the model for the 1980-99 projections.

SIMULATED WATER-LEVEL CHANGES

Sixteen model simulations were made projecting streamflow at the Colorado-Kansas State line and pumpage from January 1980 through December 1999. For discussion herein, these projections are numbered 1 through 16. Water-level changes simulated by the model from January 1, 1980, to December 31, 1999, were used to illustrate the aquifer response.

Use of cyclic patterns of streamflow at the Colorado-Kansas State line and pumpage caused the alluvial aquifer to reach steady-state conditions after a period of time ranging from 3 to 5 years. Steady state is described as a state of dynamic equilibrium in which recharge virtually balances discharge. There is no long-term water-level change nor change of water in storage. In the strictest sense a constant, or single, steady-state condition will never exist in a large-scale stream-aquifer system due to year-to-year changes in recharge and discharge (for example: pumpage, streamflow, precipitation, etc.). However, if the hydrologic system undergoes changes of a uniform and cyclic nature, such as the cyclic patterns of streamflow and pumpage used in the model projections, then the system may reach a dynamic steady-state condition where changes in storage are consistent within each cycle.

After steady-state conditions are reached, the water level in the aquifer will change in response to a change in streamflow to maintain equilibrium. The streamflow at the Colorado-Kansas State line was cycled every 1 year for option B and every 5 years for option C. Therefore, after steady state is reached, the water-level altitude in the aquifer will be similar whenever the streamflow at the State line is repeated (every 1 and 5 years). Because streamflow option A is a 10-year repeating cycle, the water-level altitude on December 31, 1999, will be similar to the altitude on December 31, 1990. Therefore, a water-level-change map from January 1980 to December 1999 is also representative of the change from January 1980 to December 1989 in each of the options.

The model boundaries on the east and west maintained the water level at the December 31, 1979, altitude throughout the 1980-99 projections. Therefore, the water-level change at these boundaries was artificially held at zero, and the nearby changes should be viewed with caution. A 1980-99 projection, using a constant rate of water-level decline equal to 0.4 ft/yr (as used in the previous study by Barker and others, 1983) and based on the observed decline of water levels between 1970-79 in wells near the east and west boundaries of the model area, was made to determine the effect of the imposed constant heads at the east and west boundaries. Pumpage option 1 and streamflow option A were used for the projection. Water levels ranged from 8 feet lower at the boundaries to less than 0.2 foot lower near the center of the model area using a constant rate of decline as compared to using constant heads. Boundary inflow was 55 percent lower, and boundary outflow was 5 percent higher using a constant rate of decline, resulting in the greater simulated water-level declines. The projection indicated that the effect of the imposition of constant heads for the east and west boundaries does not significantly affect water levels except for areas near the boundaries. The net annual river loss was 5 percent (1,228 acre-ft/yr) greater for the projection using a constant rate of decline as compared to using constant heads, partially compensating for the decreased net boundary flow available to the system experienced using a constant rate of decline. A uniform rate of water-level decline was not used for the 16 projections because of the unpredictability of water-level trends in the distant future.

In the following sections the projections using pumpage options 1 through 5 are compared for streamflow options A, B, and C. Pumpage option 6 and streamflow option D are discussed separately. Pumpage options 1, 3, and 4 are variations in the number of wells in the study area, using the continued 1979 conditions of pumpage. Pumpage options 2 and 5 are variations in the number of wells in the study area, with each well pumping the appropriated amount of water on the water right or well application. Pumpage option 6 increased the 1979 pumpage rate for existing wells by 50 percent. The projection numbers of the combinations of streamflow and pumpage rates used are shown in table 1.

Table 1.--Relationship among streamflow options, pumpage options, and projection numbers

Pumpage option	Streamflow option			
	A	B	C	D
Projection number				
1	1	6	11	NC ^{1/}
2	2	7	12	NC
3	3	8	13	NC
4	4	9	14	NC
5	5	10	15	NC
6	NC	NC	NC	16

¹ NC = not compared.

In this report, "continued 1979 pumpage rates" refers to pumpage option 1, which includes 147 wells and 60,540 acre-feet of pumpage. Although the terminology is similar to that used under "continued 1979 conditions" in the previous report by Barker and others (1983), it should be noted that none of the projections use exactly the same conditions. Pumpage option 3 in this report is similar in comparison to the number of wells and pumpage used by Barker and others (1983); however, the streamflow conditions that were used with option 3 are not similar to the conditions used by Barker and others (1983).

Streamflow Option A and Pumping Options 1-5

For streamflow option A, streamflow at the Colorado-Kansas State line was cycled on a 10-year basis using the 1971-80 gaged daily streamflow near Coolidge. This section describes the effects of pumpage options 1 through 5 and streamflow option A on the change in water levels in the aquifer and net annual river loss. Water-level-change maps for 1980-99 for each projection (figs. 4-19) were derived using 1980-99 simulated water-level changes at all nodes in the digital model grid (Barker and others, 1983).

Projection 1 combined streamflow option A and pumpage option 1. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 4). Most of the rise in water levels occurred near the northern boundary along the river. As shown in table 2, pumpage in projection 1 averaged 60,540 acre-ft/yr, deep percolation averaged 50,130 acre-ft/yr, and net river loss averaged 23,360 acre-ft/yr.

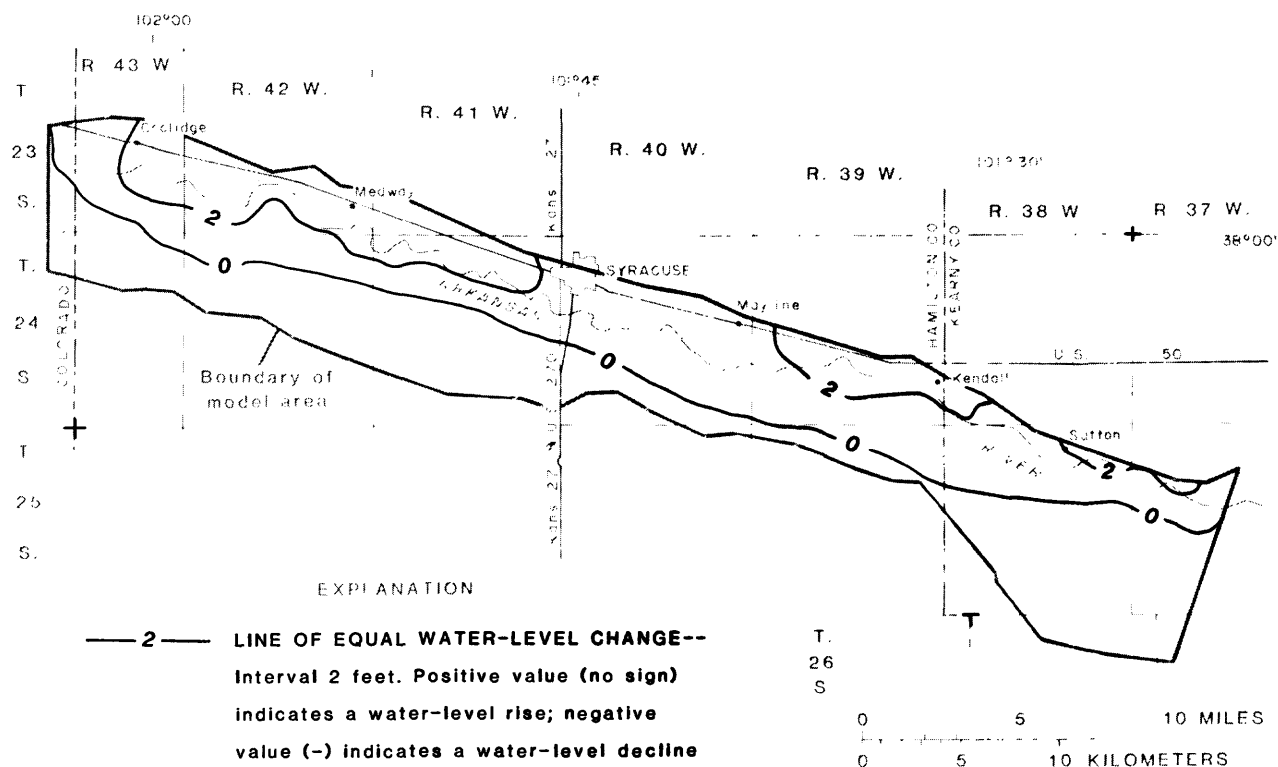


Figure 4.--Water-level change, 1980-99, using streamflow option A and pumpage option 1 (projection 1).

Table 2.--Pumpage, deep percolation, and net river loss for projections 1-16

Projection number	Pumpage option	Stream-flow option	Number of wells	Pumpage (acre-feet per year)	Deep percolation (acre-feet per year)	Net river loss (acre-feet per year)
1	1	A	147	60,540	50,130	23,360
2	2	A	1 ¹ /152	46,030	35,620	23,170
3	3	A	159	64,700	52,520	24,140
4	4	A	178	71,960	57,320	25,130
5	5	A	1 ¹ /183	57,840	43,240	24,930
6	1	B	147	60,540	50,130	26,380
7	2	B	152	46,030	35,620	26,310
8	3	B	159	64,700	52,250	27,340
9	4	B	178	71,960	57,320	28,740
10	5	B	183	57,840	43,240	28,660
11	1	C	147	60,540	48,500	21,450
12	2	C	152	46,030	33,940	21,150
13	3	C	159	64,700	50,900	22,050
14	4	C	178	71,960	55,610	22,610
15	5	C	183	57,840	41,500	22,280
16	6	2/D	159	97,050	80,280	11,140

¹ Pumpage options 2 and 5 include five wells that did not pump during 1978 or 1979.

² Projection 16 uses continued 1979 streamflow.

Projection 2 combined streamflow option A and pumpage option 2. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 5). Most of the rise in water levels occurred near the northern boundary along the river. Pumpage in projection 2 averaged 46,030 acre-ft/yr, deep percolation averaged 35,620 acre-ft/yr, and net river loss averaged 23,170 acre-ft/yr.

Comparing projection 2 with projection 1 revealed that pumping a quantity of water equal to the amounts appropriated on water rights, instead of continuing 1979 pumpage rates, would decrease the annual pumpage by 24 percent. However, net annual river loss would decrease by only 1 percent. The water-level-change maps appear similar (figs. 4 and 5).

The ratio of deep percolation to pumpage was higher (83 percent) when wells pumped at the continued 1979 rate rather than pumping the appropriated amount (77 percent). In this case, increased pumpage from a given number of wells had the effect of recycling the additional water by recharge back to the aquifer, resulting in a negligible difference in the amount of ground water in storage.

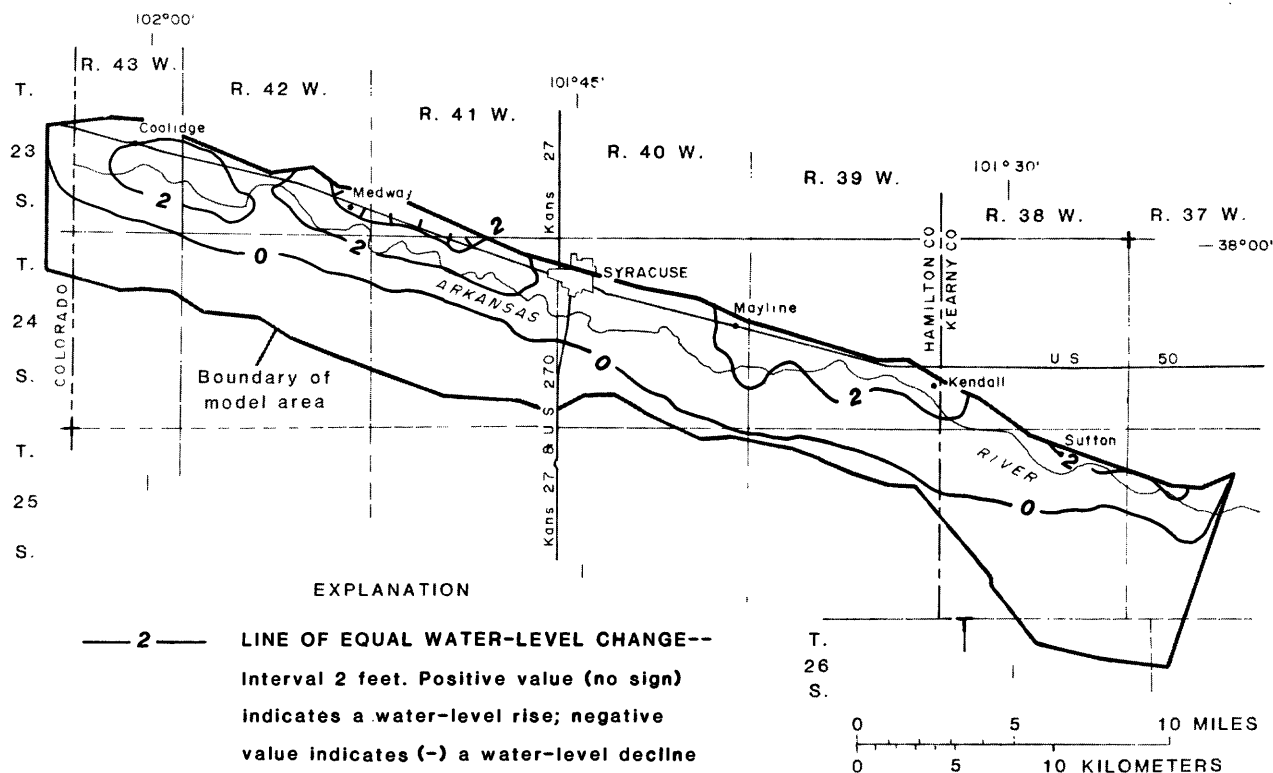


Figure 5.--Water-level change, 1980-99, using streamflow option A and pumpage option 2 (projection 2).

Projection 3 combined streamflow option A with pumpage option 3. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 6). Most of the rise in water levels occurred near the northern boundary along the river. Pumpage in projection 3 averaged 64,700 acre-ft/yr, deep percolation averaged 52,520 acre-ft/yr, and net river loss averaged 24,140 acre-ft/yr.

Comparing projection 3 with projection 1 revealed that including pumpage from existing wells for which permits were pending increased the annual pumpage by 7 percent and increased the net annual river loss by 3 percent. The water-level-change maps are very similar (figs. 4 and 6).

When pumpage was increased over the 1979 rates in the modeled area and distributed to new pumpage locations, additional ground water was removed from storage in the aquifer and net annual river loss increased. This is because the additional crops irrigated at the new sites increased the crop consumptive use. Therefore, more water was consumed by plants in the soil zone, and the ratio of deep percolation to pumpage decreased to 81 percent. The increased pumpage was not recycled completely back to the aquifer as it was in projection 1.

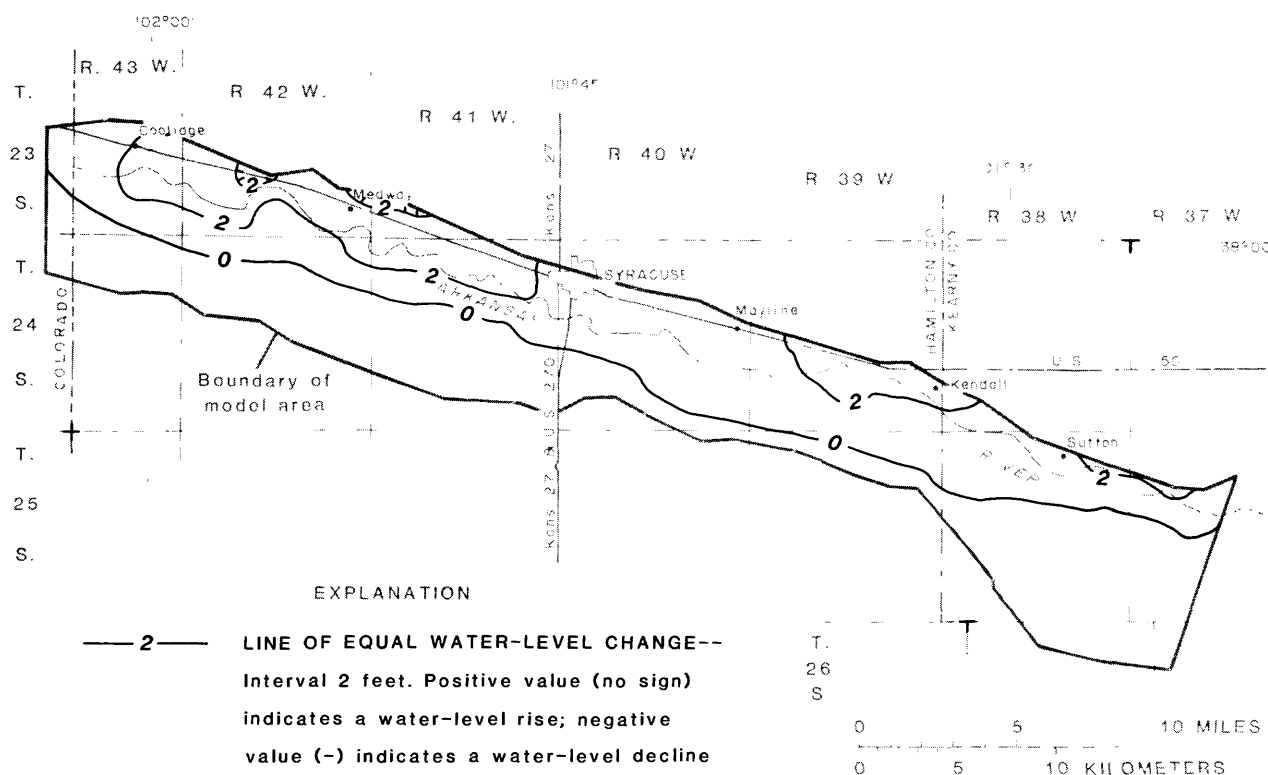


Figure 6.--Water-level change, 1980-99, using streamflow option A and pumpage option 3 (projection 3).

Projection 4 combined streamflow option A with pumpage option 4. The water-level change ranged from less than 4 feet of decline to less than 4 feet of rise (fig. 7). Most of the rise in water levels occurred near the northern boundary along the river. Pumpage in projection 4 averaged 71,960 acre-ft/yr, deep percolation averaged 57,320 acre-ft/yr, and net river loss averaged 25,130 acre-ft/yr.

Comparing projection 4 with projection 1 revealed that including both existing and proposed wells for which permits were pending increased the annual pumpage by 19 percent and increased the net annual river loss by 8 percent. A comparison of the water-level-change maps indicated that the rise in water levels would be slightly less in projection 4 than in projection 1 (figs. 4 and 7).

Projection 5 combined streamflow option A with pumpage option 5. The water-level change ranged from less than 4 feet of decline to less than 4 feet of rise (fig. 8). The largest area of decline occurred in the southwestern part of the model area. Pumpage in projection 5 averaged 57,840 acre-ft/yr, deep percolation averaged 43,240 acre-ft/yr, and net river loss averaged 24,930 acre-ft/yr.

Comparing projection 5 with projection 2 revealed that including both existing and proposed wells for which permits were pending increased the appropriated annual pumpage by 26 percent and increased the net annual river loss by 8 percent. A comparison of the water-level-change maps indicated that the rise in water levels would be noticeably less in projection 5 than in projection 2 (figs. 5 and 8), which explains the increase in the net annual river loss.

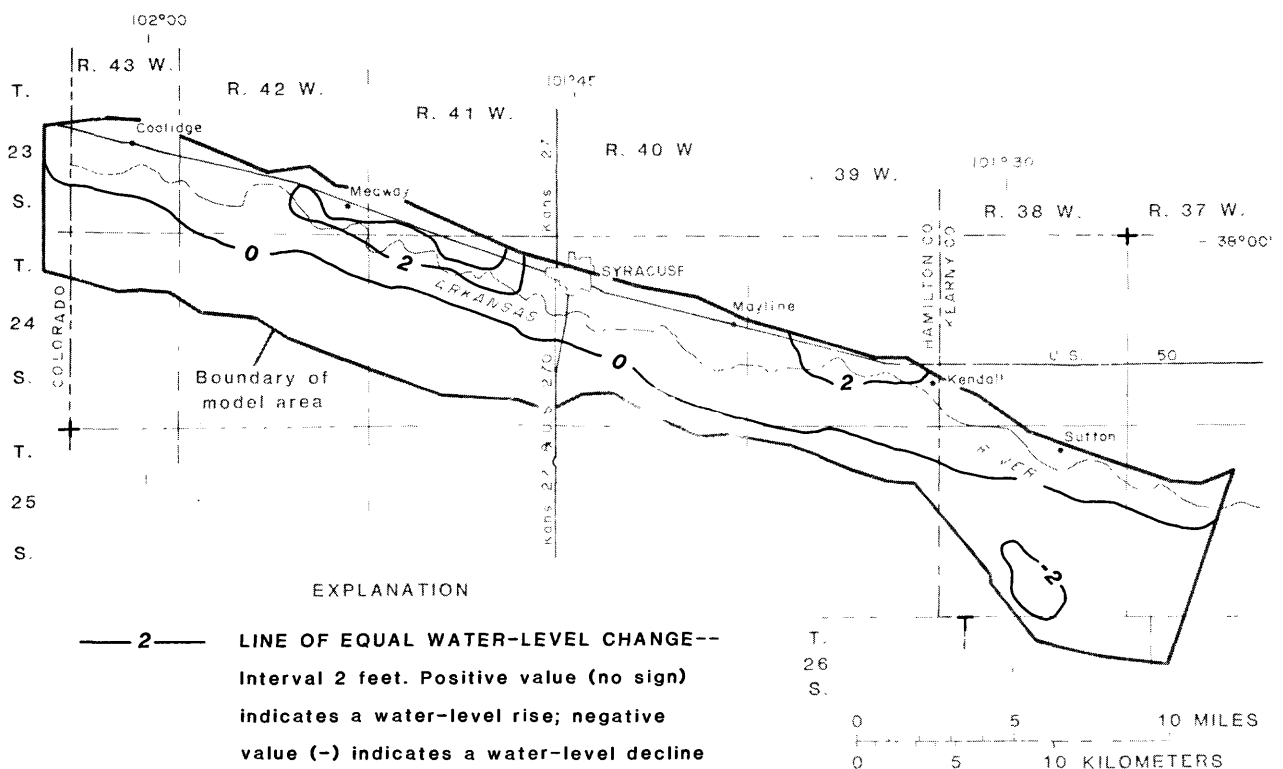


Figure 7.--Water-level change, 1980-99, using streamflow option A and pumpage option 4 (projection 4).

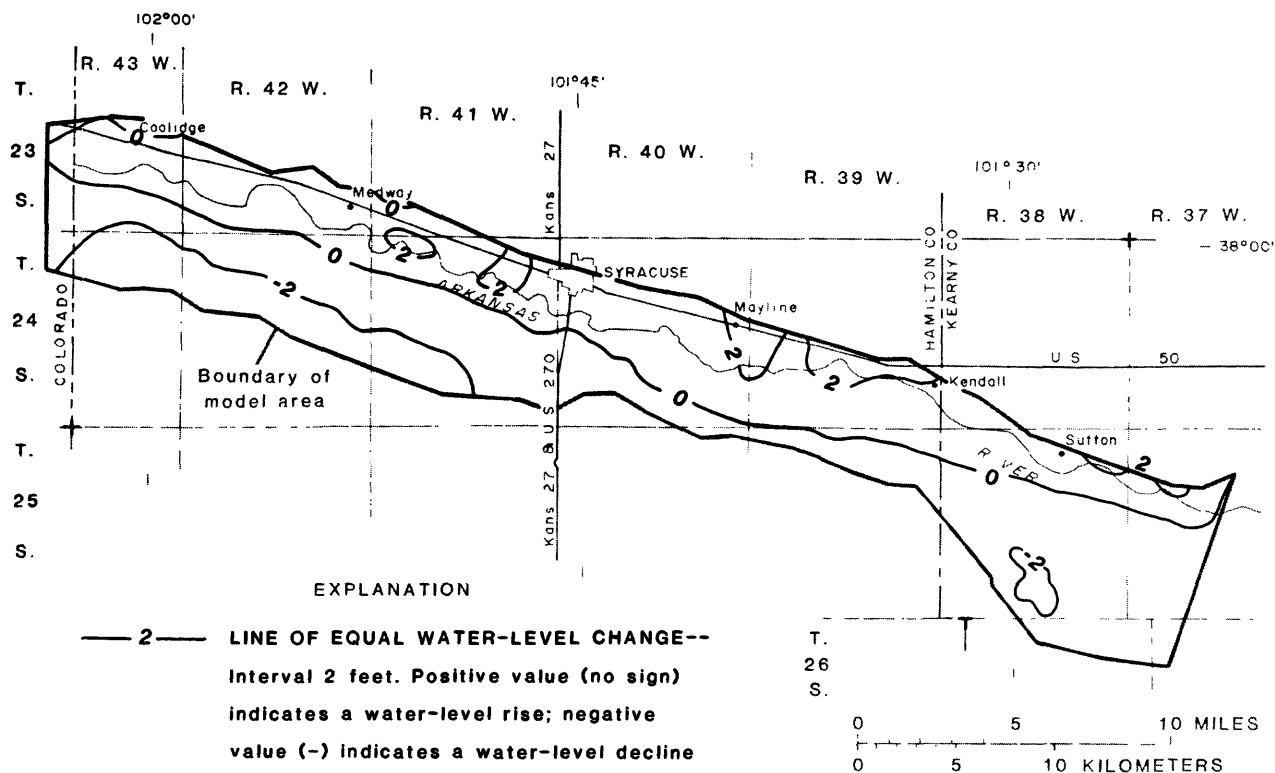


Figure 8.--Water-level change, 1980-99, using streamflow option A and pumpage option 5 (projection 5).

The results of a comparison between projection 5 and projection 4 are essentially the same as the comparison between projection 2 and projection 1. There is a greater difference in the water-level-change maps for projections 4 and 5 because there is a greater difference in ground-water storage than there is between projections 1 and 2. Much of the additional water from increased pumpage again was recycled back to the aquifer.

Streamflow Option B and Pumpage Options 1-5

For streamflow option B, the Colorado-Kansas State line streamflow of the Arkansas River was cycled on an annual basis using the gaged 1971-80 average monthly streamflow near Coolidge. This section describes the effects of pumpage options 1 through 5 and streamflow option B on the changes in water levels and net annual river loss.

Projection 6 combined streamflow option B and pumpage option 1. The water-level change ranged from less than 2 feet of decline to less than 6 feet of rise (fig. 9). Most of the rise occurred along the river in the western one-half of the study area. Pumpage in projection 6 averaged 60,540 acre-ft/yr, deep percolation averaged 50,130 acre-ft/yr, and net river loss averaged 26,380 acre-ft/yr.

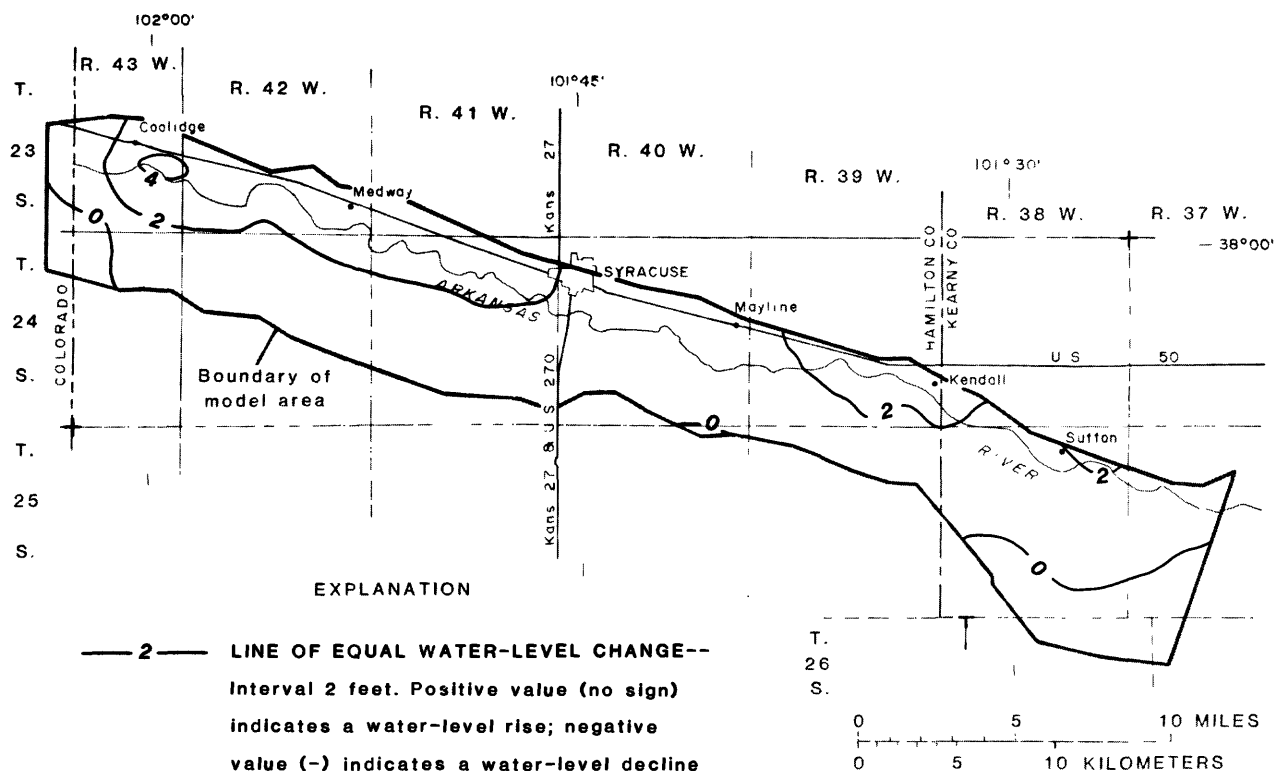


Figure 9.--Water-level change, 1980-99, using streamflow option B and pumpage option 1 (projection 6).

Projection 7 combined streamflow option B and pumpage option 2. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 10). Most of the rise occurred along the river in the western one-half of the modeled area. Pumpage in projection 7 averaged 46,030 acre-ft/yr, deep percolation averaged 35,620 acre-ft/yr, and net river loss averaged 26,310 acre-ft/yr.

Comparing projection 7 with projection 6 revealed that if the amount appropriated by water rights was pumped, then net annual river loss would decrease by less than 1 percent. The water-level-change maps appear similar (figs. 9 and 10). This is because a higher percentage of the pumped water in projection 6 (83 percent) returns to the aquifer via deep percolation than in projection 7 (77 percent). In this case, increased pumpage from a given number of wells had the effect of recycling the additional water back to the aquifer, resulting in a negligible difference in the ground water in storage and the net annual river loss.

Projection 8 combined streamflow option B with pumpage option 3. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 11). Most of the rise occurred along the river in the western one-half of the model area. Pumpage in projection 8 averaged 64,700 acre-ft/yr, deep percolation averaged 52,520 acre-ft/yr, and net river loss averaged 27,340 acre-ft/yr.

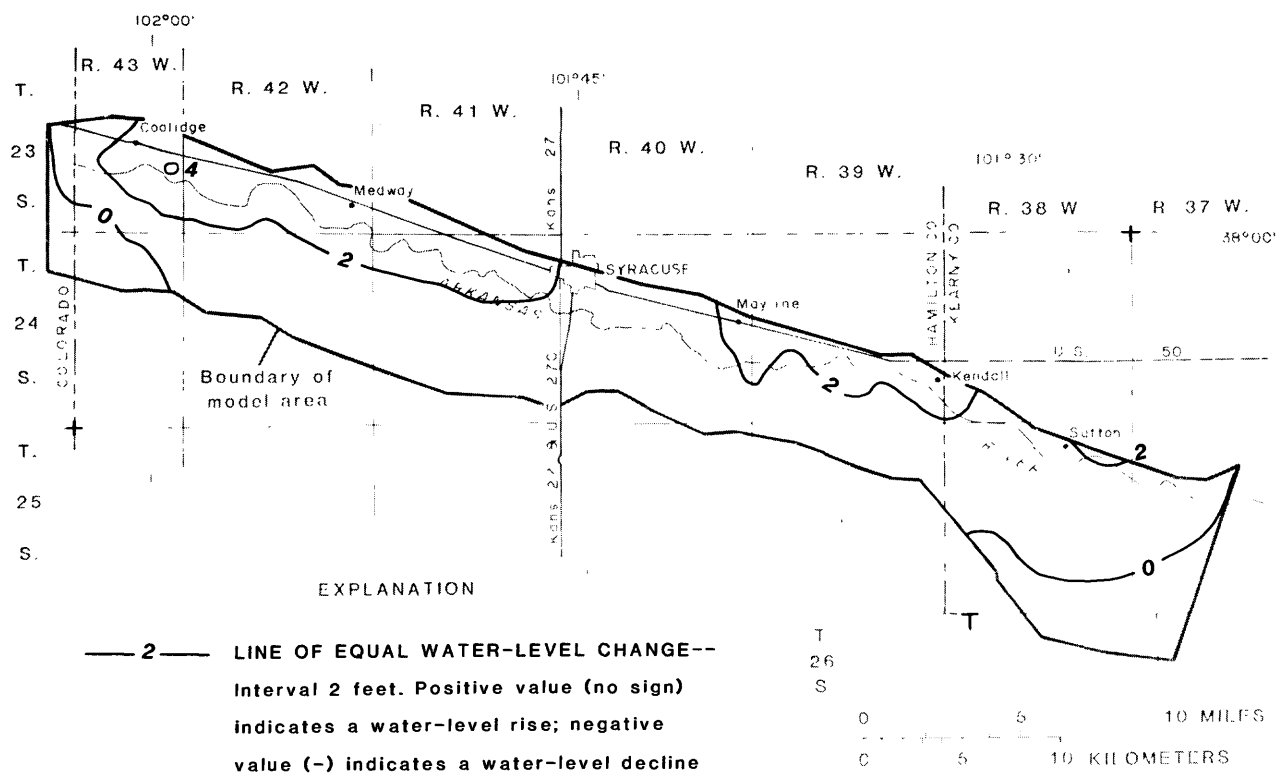


Figure 10.--Water-level change, 1980-99, using streamflow option B and pumpage option 2 (projection 7).

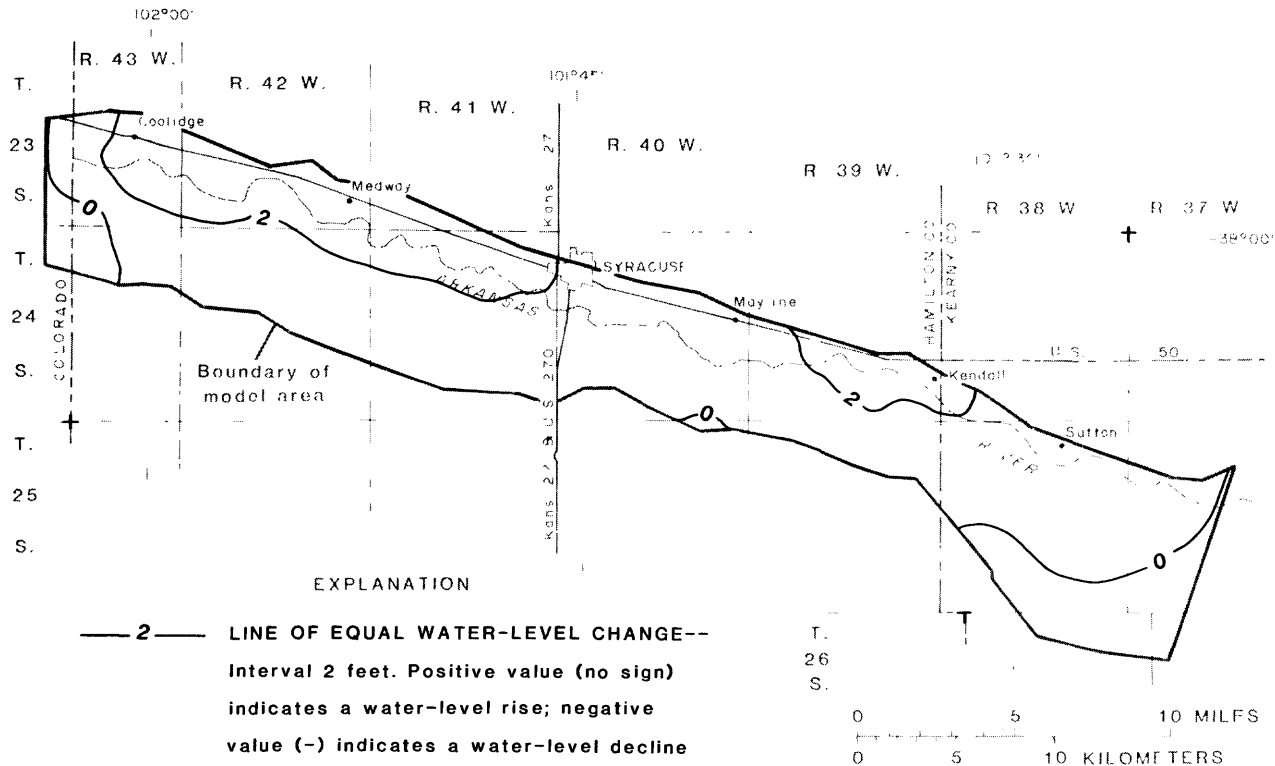


Figure 11.--Water-level change, 1980-99, using streamflow option B and pumpage option 3 (projection 8).

Comparing projection 8 with projection 6 revealed that including the existing wells for which permits were pending increased the net annual river loss by 4 percent. The water-level-change maps are similar (figs. 9 and 11). As in streamflow option A, when pumpage in the model area was increased in the form of new well sites, additional ground water was removed from storage in the aquifer and net annual river loss increased. This is because when additional well sites were included in the study area, crop consumptive use was increased with the additional crops, and the ratio of deep percolation to pumpage decreased. Therefore, the increased pumpage was not completely recycled back to the aquifer as it was in projection 6.

Projection 9 combined streamflow option B with pumpage option 4. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 12). Most of the rise was along the river in the western one-half of the study area. Pumpage in projection 9 averaged 71,960 acre-ft/yr, deep percolation averaged 57,320 acre-ft/yr, and net river loss averaged 28,740 acre-ft/yr.

Comparing projection 9 with projection 6 revealed that including both existing and proposed wells for which permits were pending increased the net annual river loss by 9 percent. A comparison of the water-level-change maps indicated that the rise in water levels would be slightly less in projection 9 than in projection 6 (figs. 9 and 12), which explains the increase in the net annual river loss.

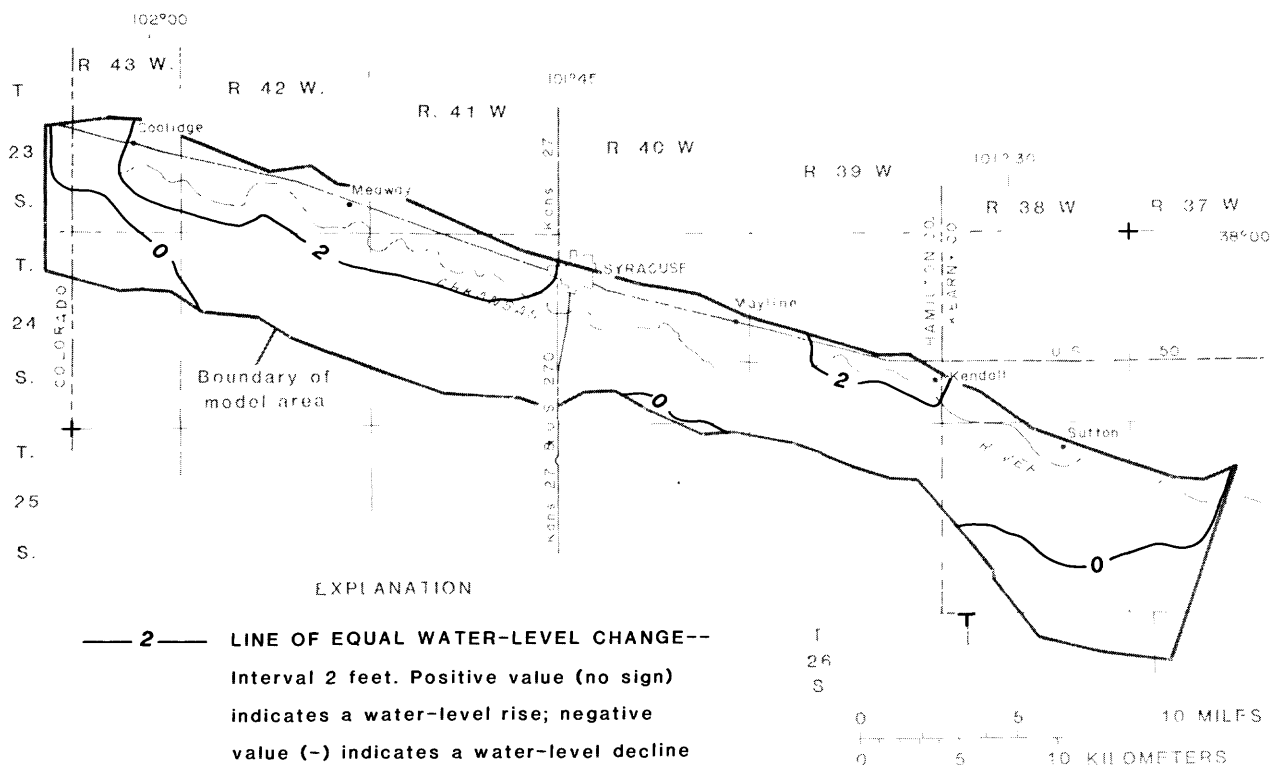


Figure 12.--Water-level change, 1980-99, using streamflow option B and pumpage option 4 (projection 9).

Projection 10 combined streamflow option B with pumpage option 5. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 13). The largest area of rise occurred along the river in the western one-half of the study area. Pumpage in projection 10 averaged 57,840 acre-ft/yr, deep percolation averaged 43,240 acre-ft/yr, and net river loss averaged 28,660 acre-ft/yr.

Comparing projection 10 with projection 7 revealed that including both existing and proposed wells for which permits were pending increased the net annual river loss by 9 percent. A comparison of the water-level-change maps indicated that the rise in water levels would be slightly less in projection 10 than in projection 7 (figs. 10 and 13), which explains the increase in the net annual river loss.

Streamflow Option C and Pumpage Options 1-5

For streamflow option C, the Colorado-Kansas State line streamflow of the Arkansas River was cycled on a 5-year basis using the 1976-80 gaged daily streamflow near Coolidge. This section describes the effects of pumpage options 1 through 5 and streamflow option C on the change in water levels and net annual river loss.

Projection 11 combined streamflow option C and pumpage option 1. The water-level change ranged from less than 2 feet of decline to less than 4 feet of rise (fig. 14). Most of the decline occurred along the southern boundary. Pumpage in projection 11 averaged 60,540 acre-ft/yr, deep percolation averaged 48,500 acre-ft/yr, and net river loss averaged 21,450 acre-ft/yr.

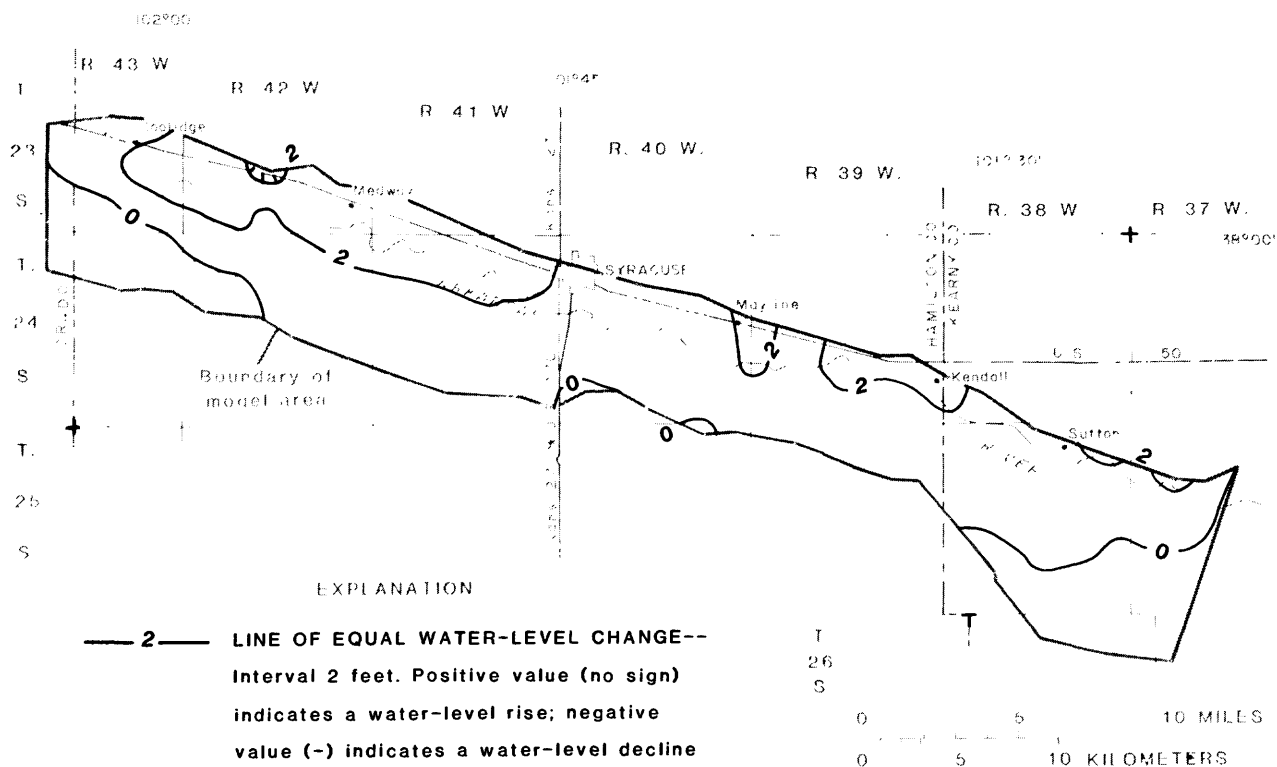


Figure 13.--Water-level change, 1980-99, using streamflow option B and pumpage option 5 (projection 10).

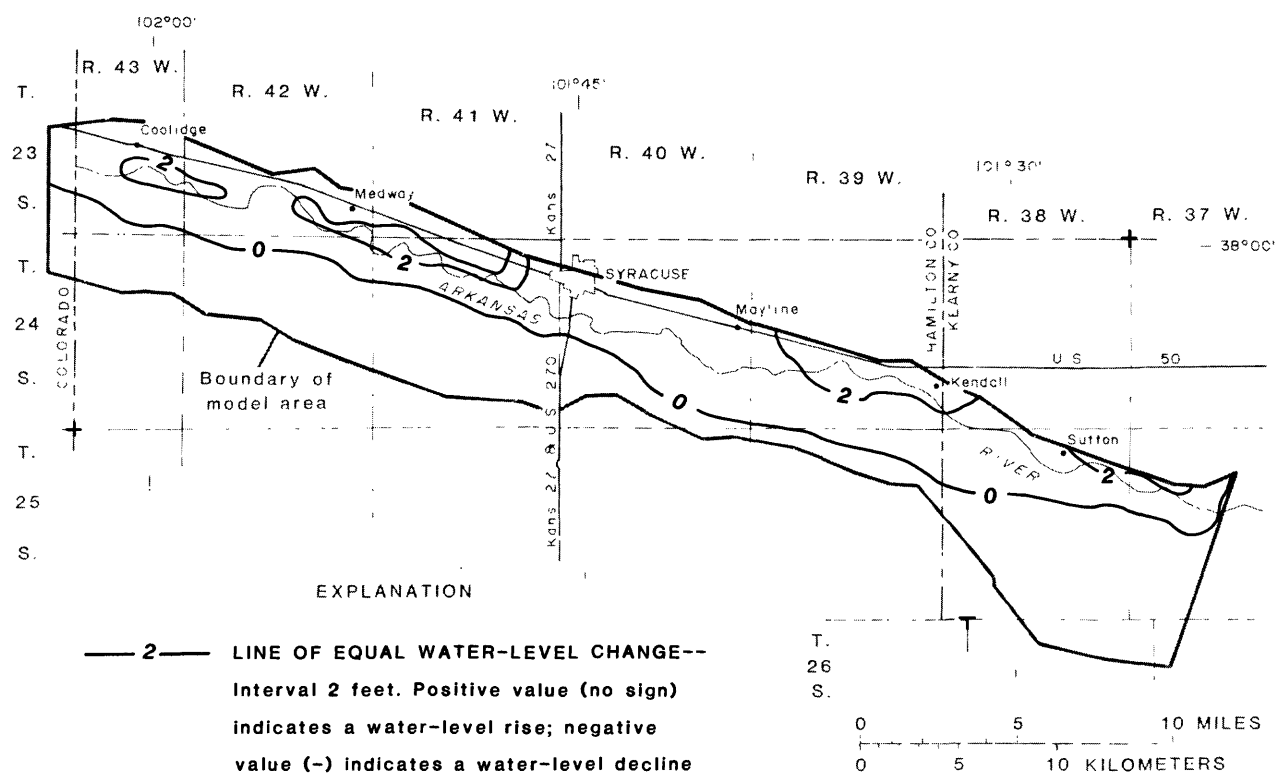


Figure 14.--Water-level change, 1980-99, using streamflow option C and pumpage option 1 (projection 11).

Projection 12 combined streamflow option C and pumpage option 2. The water-level change ranged from less than 4 feet of decline to less than 4 feet of rise (fig. 15). Most of the decline occurred along the southern boundary of the modeled area. Pumpage in projection 12 averaged 46,030 acre-ft/yr, deep percolation averaged 33,940 acre-ft/yr, and net river loss averaged 21,150 acre-ft/yr.

Comparing projection 12 with projection 11 revealed that if the amount appropriated by water rights was pumped, then the net annual river loss would decrease by 1 percent. The effect of increased pumpage in projection 11 was negated by increased deep percolation.

Projection 13 combined streamflow option C with pumpage option 3. The water-level change ranged from less than 4 feet of decline to less than 4 feet of rise (fig. 16). Most of the decline was in the southern one-half of the modeled area. Pumpage in projection 13 averaged 64,700 acre-ft/yr, deep percolation averaged 50,900 acre-ft/yr, and net river loss averaged 22,050 acre-ft/yr.

Comparing projection 13 with projection 11 revealed that adding the existing wells for which permits were pending increased the net annual river loss by 3 percent. A comparison of the water-level-change maps indicated that the rise in water levels would be slightly less in projection 13 than in projection 11 (figs. 14 and 16), which explains the increase in the net annual river loss.

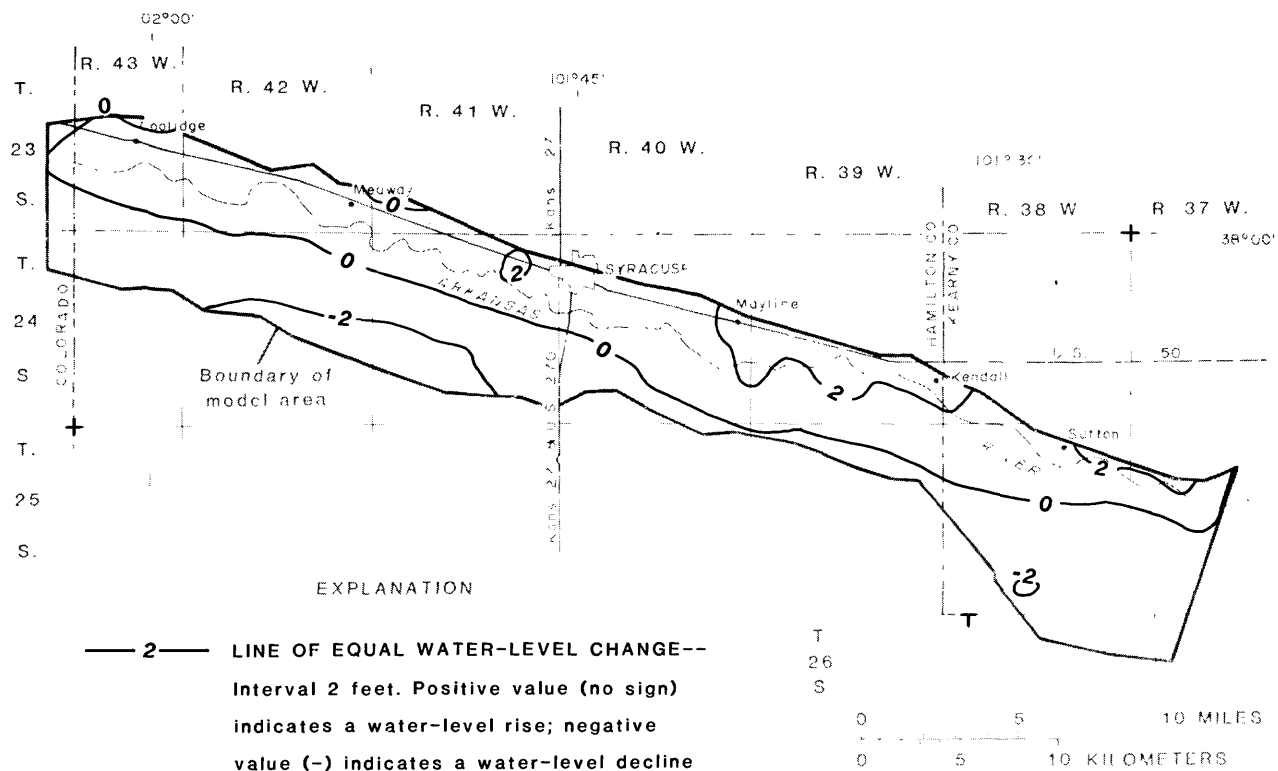


Figure 15.--Water-level change, 1980-99, using streamflow option C and pumpage option 2 (projection 12).

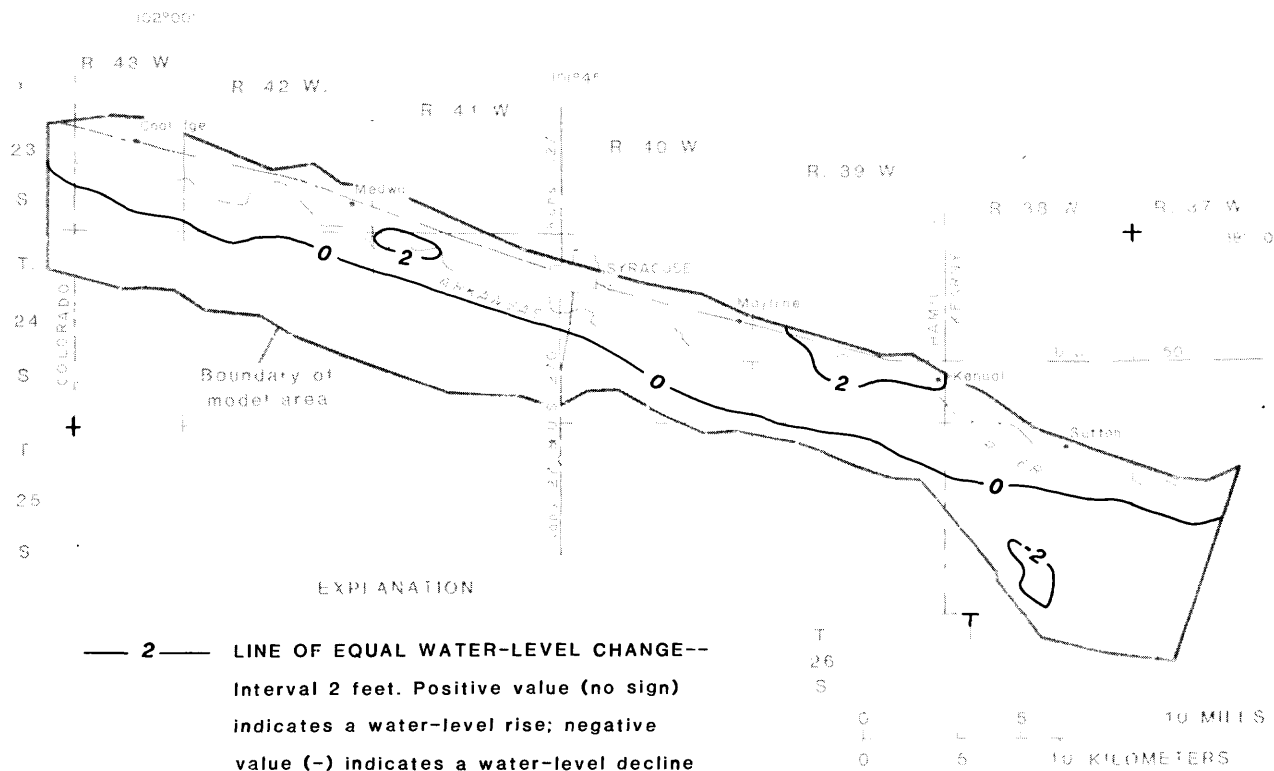


Figure 16.--Water-level change, 1980-99, using streamflow option C and pumpage option 3 (projection 13).

Projection 14 combined streamflow option C with pumpage option 4. The water-level change ranged from less than 4 feet of decline to less than 4 feet of rise (fig. 17). The largest area of decline was in the southwestern part of the model area. Pumpage in projection 14 averaged 71,960 acre-ft/yr, deep percolation averaged 55,610 acre-ft/yr, and net river loss averaged 22,610 acre-ft/yr.

Comparing projection 14 with projection 11 revealed that including both existing and proposed wells for which permits were pending increased the net annual river loss by 5 percent. A comparison of the water-level-change maps indicated that the rise in water levels would be noticeably less in projection 14 than in projection 11 (figs. 14 and 17), which explains the increased net annual river loss.

Projection 15 combined streamflow option C with pumpage option 5. The water-level change ranged from less than 4 feet of decline to less than 4 feet of rise (fig. 18). Most of the decline was in the southwestern part of the study area. Pumpage in projection 15 averaged 57,840 acre-ft/yr, deep percolation averaged 41,500 acre-ft/yr, and net annual river loss averaged 22,280 acre-ft/yr.

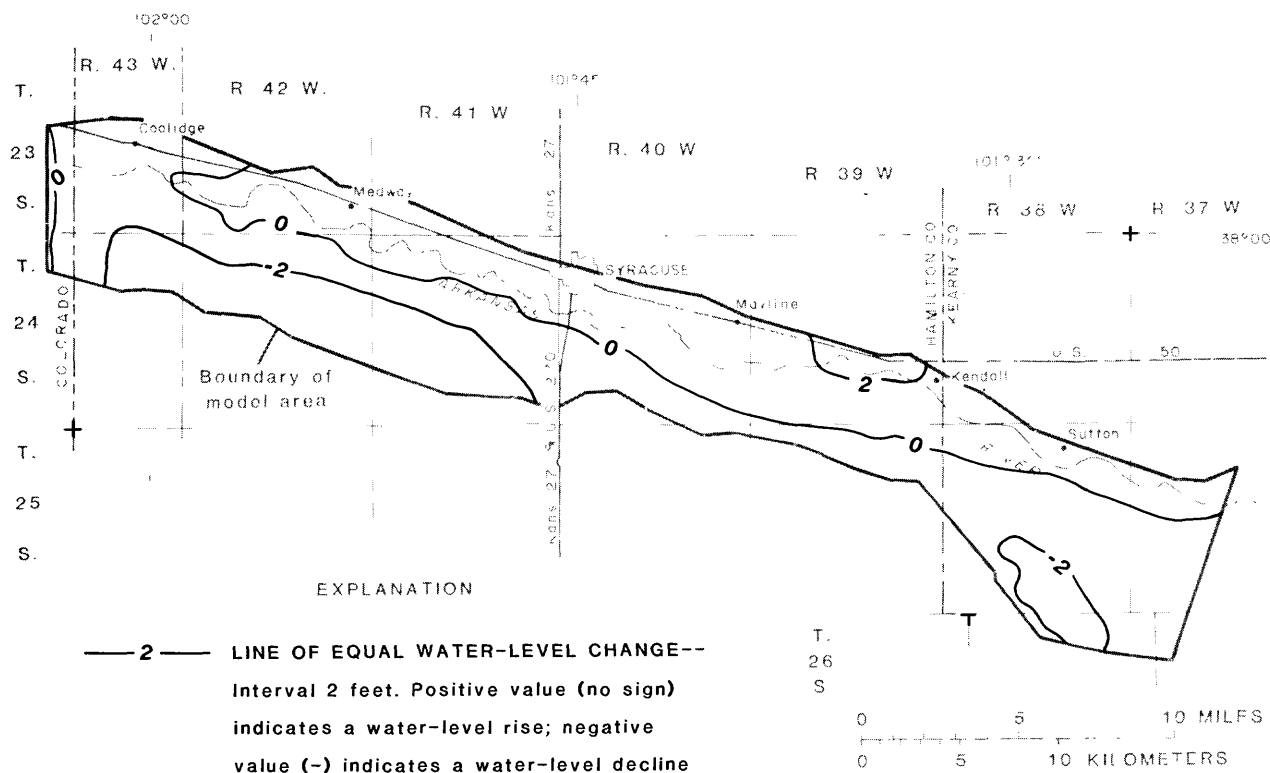


Figure 17.--Water-level change, 1980-99, using streamflow option C and pumpage option 4 (projection 14).

Comparing projection 15 with projection 12 revealed that including both existing and proposed wells for which permits were pending increased the net annual river loss by 5 percent. A comparison of the water-level-change maps indicated that the water-level declines were greater and the water-level rises were less in projection 15 than in projection 11 (figs. 15 and 18), which explains the increased net annual river loss.

Net annual river losses were smaller for streamflow option C than for options A and B because 1976-80 streamflow was generally less than 1971-75 streamflow. Deep percolation was less for streamflow option C than for options A and B because 1976-80 surface-water diversions were also less than those for 1971-75.

Streamflow Option D and Pumpage Option 6

An additional projection was made by cycling the 1979 gaged daily streamflow of the Arkansas River near Coolidge for 1980-99. The 1979 gaged flow was the lowest streamflow of the 1970's. The pumpage rate for wells that pumped during 1979 (pumpage option 3) was increased by 50 percent (pumpage option 6).

The water-level change for projection 16 ranged from 0 (resulting from the imposed constant heads at the east and west boundaries of the modeled area) to less than 8 feet of decline (fig. 19). Pumpage in projection 16 averaged 97,050 acre-ft/yr, deep percolation averaged 80,280 acre-ft/yr, and net river loss averaged 11,140 acre-ft/yr. This projection showed that water levels in the modeled area would have continued to decline in the next 20 years if the low streamflow that occurred during 1979 had continued and ground-water pumpage had increased. The net annual river loss was smaller because of the decreased streamflow in the river.

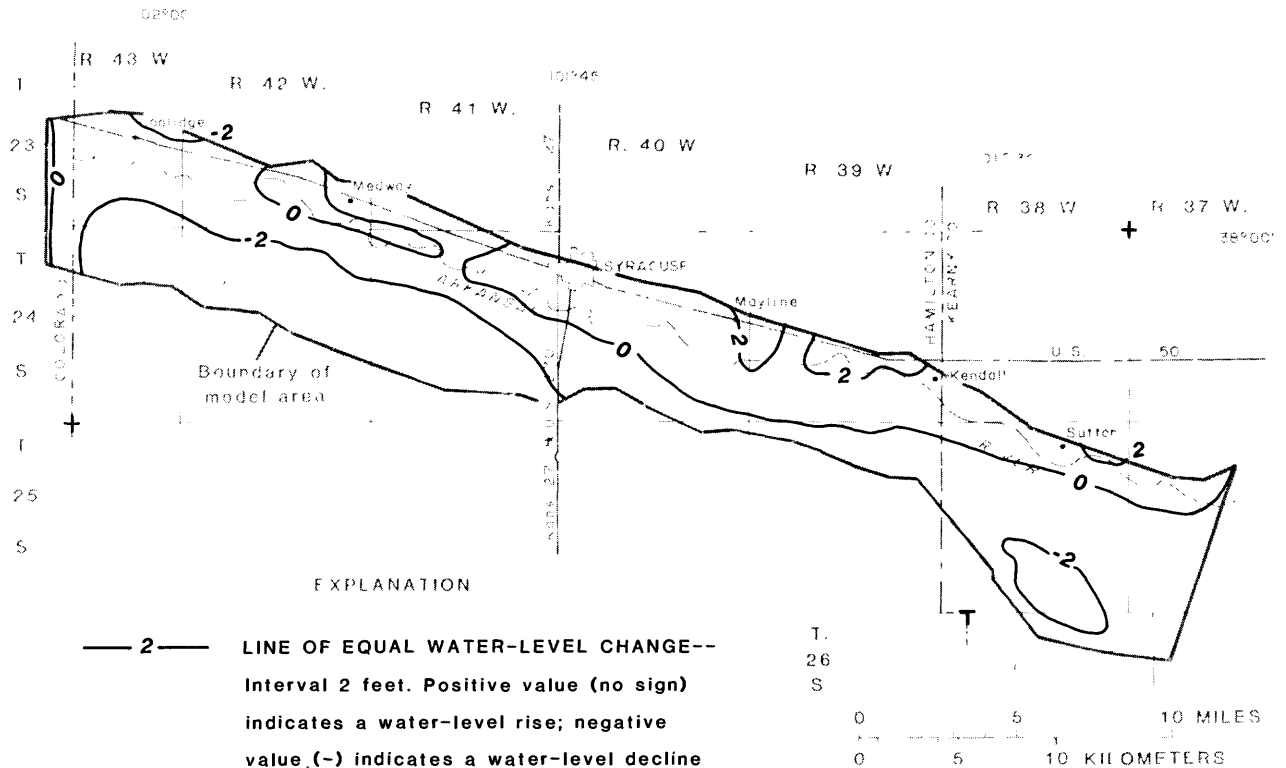


Figure 18.--Water-level change, 1980-99, using streamflow option C and pumpage option 5 (projection 15).

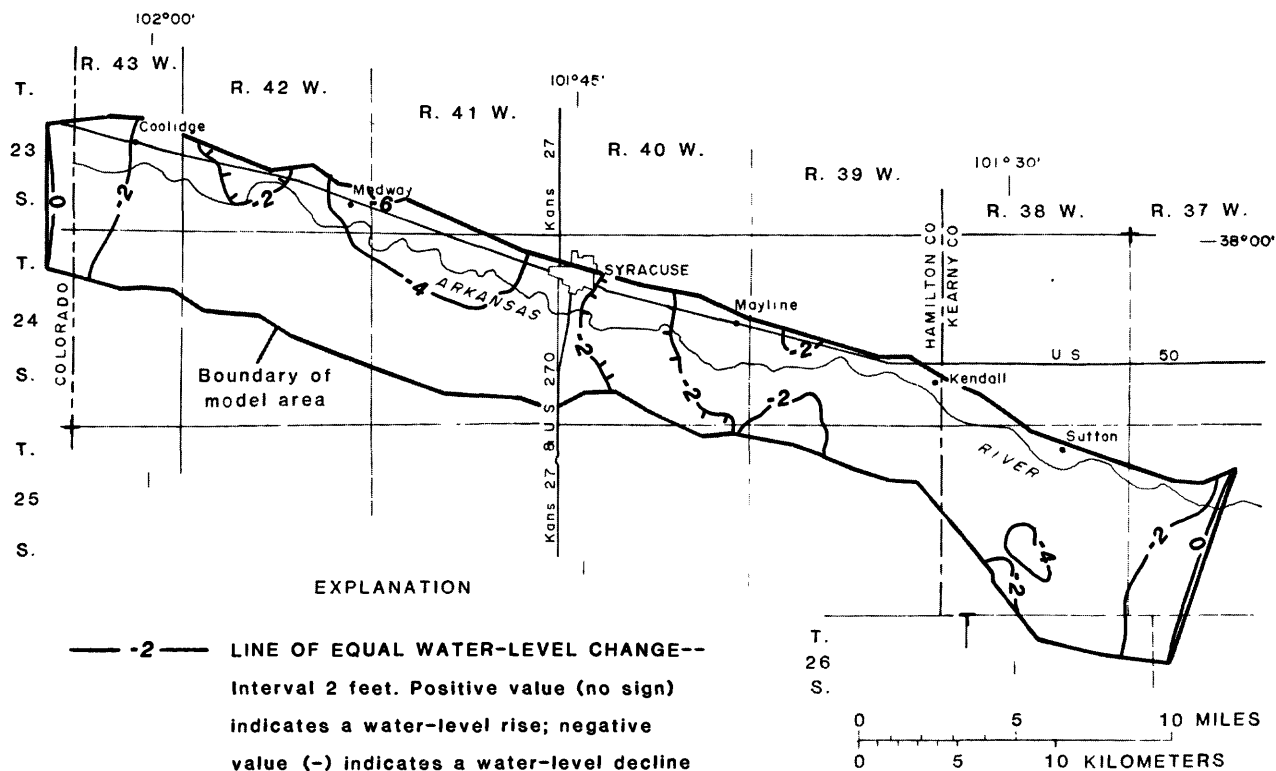


Figure 19.--Water-level change, 1980-99, using streamflow option D and pumpage option 6 (projection 16).

SUMMARY

Information was needed by the Kansas State Board of Agriculture, Division of Water Resources, on the effects that additional ground-water development would have on streamflow and water levels in an area along the Arkansas River in Hamilton and Kearny Counties, southwestern Kansas. A computer model, based on a finite-element numerical technique, was used to simulate the changes in streamflow and water levels from 1980 through 1999. Four hypothetical streamflow options and six hypothetical pumpage options were tested in the model. Pumpage options were variations in the number of wells using either the continued 1979 conditions of pumpage or the amount of water appropriated or requested by water-right applications. One pumpage option increased 1979 pumpage by 50 percent.

Results from 1980-99 simulations indicated that increased pumpage from a given number of wells, with no increase in irrigated acreage, has the effect of recycling the additional water back to the aquifer. If the amount appropriated by water rights was pumped rather than continued 1979 conditions of pumpage, annual pumpage in the modeled area would be reduced 24 percent. However, this reduction in pumpage would reduce the net annual river loss by only 1 percent. The ratio of deep percolation to pumpage was higher when wells continued to pump at 1979 rates rather than pumping the appropriated amounts, resulting in a negligible difference in the amount of ground water in storage and net annual river losses.

When pumpage in the modeled area was increased by adding new wells, additional ground water was removed from storage in the aquifer and net annual river loss increased. Additional crops irrigated by the new wells increase the crop consumptive use causing less water to percolate through the soil zone. Therefore, more water is consumed in the soil zone by plants, and the ratio of deep percolation to pumpage decreases. When pumpage from existing wells for which permits were pending was included with pumpage from wells with water rights, the pumpage at continued 1979 rates increased by 7 percent and the net annual river loss increased 3 to 4 percent.

When pumpage from both existing and proposed wells for which permits were pending was included with pumpage from wells with water rights, the pumpage at continued 1979 rates increased 19 percent and the net annual river loss increased 5 to 9 percent.

When both existing and proposed wells for which permits were pending and wells with water rights pumped the amounts appropriated on the water-right or well application, pumpage increased 26 percent over the pumpage from wells with water rights only. The net annual river loss increased by 5 to 9 percent.

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MODEL DOCUMENTATION

By Jerry E. Carr

The digital-computer model used to make the long-term projections for the Arkansas River valley in Hamilton and Kearny Counties was written by J. V. Tracy, U.S. Geological Survey. The Tracy finite-element model was selected because it offered a practical combination of features necessary to meet the project objectives. The model was desirable because it (1) added flexibility in grid design, (2) allowed simulation of the boundary conditions associated with the stream-aquifer interface and the bedrock-bounded limits of the relatively narrow alluvial aquifer, and (3) accounted for transient conditions of streamflow.

Included in this attachment are: a description of the model, sample problems, a description of program routines, flow charts, model-input setup, a description of program common blocks, variable definitions, and a program listing. The description of model-input setup, program routines, flow charts, and variable definitions were modified from those supplied by E. J. Wexler (U.S. Geological Survey, written commun., 1983).

Description of Model

The Tracy model is a two-dimensional, finite-element model based on the Galerkin finite-element method, described by Desai and Abel (1972). Additional descriptions of this method can be found in Pinder and Gray (1977). The model uses a direct-solution method to solve the partial-differential equations describing nonsteady, two-dimensional ground-water flow. The finite-element method involves subdividing the area to be modeled into subregions (called elements) having simple shapes and assuming that the solution to the partial-differential equation can be written as a linear combination of simple basis functions. See Desai and Abel (1972, pages 89-93, 114-117) and Pinder and Gray (1977, pages 99-101 and 115). By selecting the basis functions to be nonzero only over elements on which they are located, the computations are simplified. Associated with the basis functions are undetermined coefficients that are determined by applying the Galerkin scheme. This process results in a system of simultaneous equations where the unknowns are heads at points (called nodes) that are associated with the elements. This system of equations is solved using selected direct methods of matrix solutions.

The model is formulated to produce an approximate solution to the following partial-differential ground-water-flow equation:

$$\frac{\partial}{\partial x} \left(Kb \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial y} \left(Kb \frac{\partial h}{\partial y} \right) = S \frac{\partial h}{\partial t} + Q \quad (1)$$

where

x and y are the coordinate axes [L];

K is the hydraulic conductivity [L/t⁻¹];

b is the saturated thickness [L];

h is the hydraulic head [L];

S is the storage coefficient [dimensionless];

t is the time, [t]; and

Q = Q (x, y, t) = net vertical flux into the aquifer from point or distributed sources (sinks), such as wells, evapotranspiration, ground-water percolation, or river-aquifer interaction [L/t].

The model can be setup with the following options: steady-state or transient flow, water-table or confined aquifers, quadrangle or triangular elements, latitude-longitude or x, y coordinates, with or without stream-aquifer interaction, with or without drawdown calculations, with or without plot routines, and with or without evapotranspiration. The user controls the selection of these modeling options by defining eight true or false logical flags at the beginning of the data-input deck. The model is setup to calculate the location of interior nodes given the coordinates of the top and bottom grid nodes. An example of node arrangement and element numbering for a 20 x 20 node, triangular-element grid is shown in figure 20. Presently, the model has no option to allow the user to specify interior-node locations. The model can utilize any consistent set of length units; however, the unit of time is restricted to seconds, and precipitation, evapotranspiration, and soil capacity are entered in appropriate units of inches and months.

Boundary conditions in the model are specified for each node in the model as either interior, specified flux, or as specified head by the IBNODE array. Once defined in the IBNODE array, the specific flux or head values are given in the BQH array.

Stream-aquifer interaction is simulated in the model according to:

$$Q = \frac{K'}{b'} (h_s - h_a) A, \quad h_a \geq \text{SBA} \quad (2a)$$

or

$$\frac{K'}{b'} (h_s - \text{SBA}) A, \quad h_a < \text{SBA} \quad (2b)$$

where

Q = rate of leakage [L^3/t];

$\frac{K'}{b'}$ = streambed leakage, or ratio of hydraulic conductivity of streambed [t^{-1}];

h_s = altitude of stream stage [L];

h_a = altitude of potentiometric surface in aquifer [L]; and

A = area of streambed reach [L^2]; and

SBA = streambed altitude.

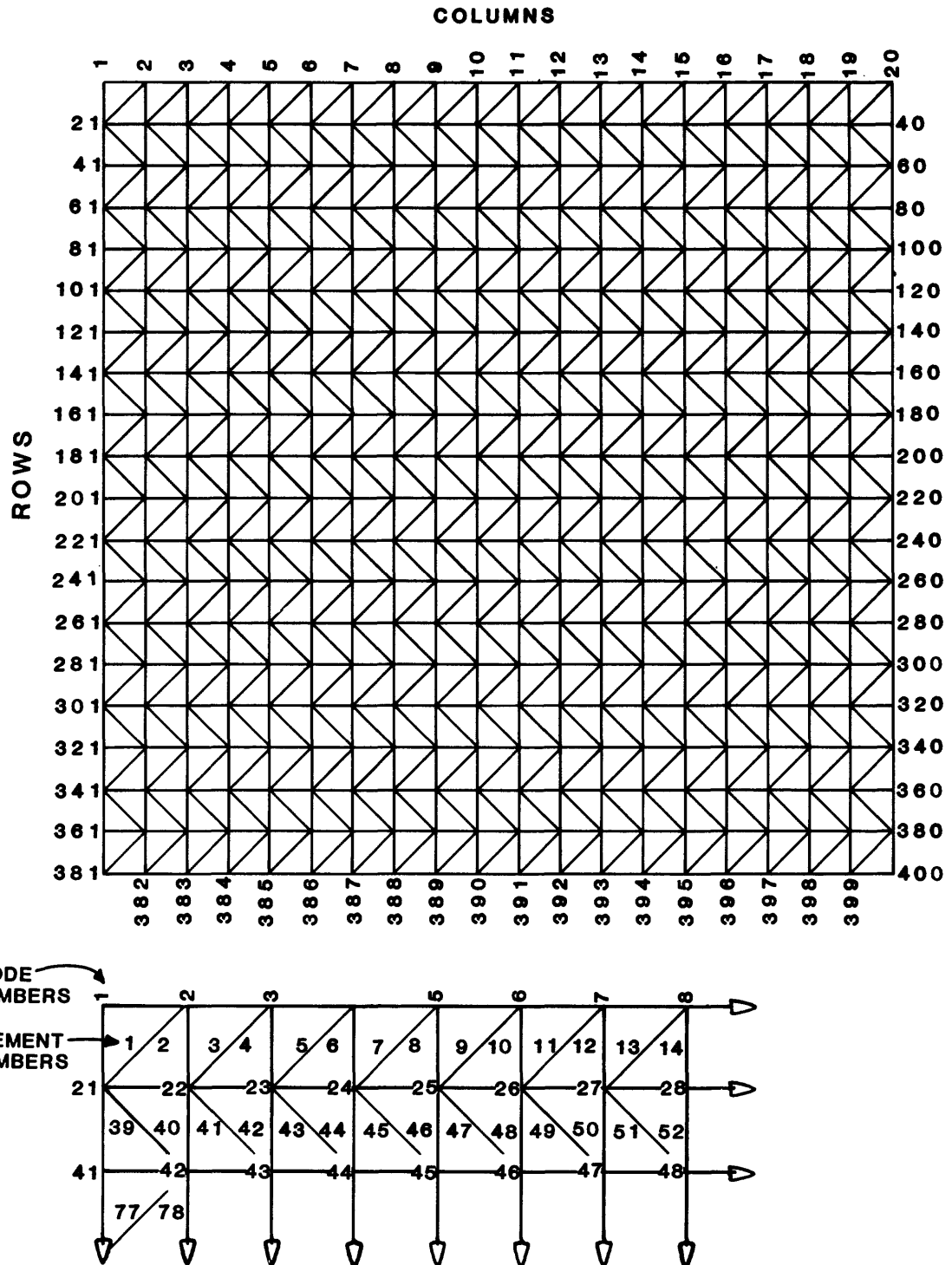


Figure 20.--Triangular nodal arrangement for finite-element grid and numbering of nodes and elements.

Streambed altitude, lengths, widths, reaches, and zones are specified for each riverbed node as part of the model input (fig. 21). The riverbed nodes are positioned such that the general path of the river channel is approximated by straight lines that link adjacent riverbed nodes. Streambed leakance is assigned by river zones, which may include more than one river node. The model simulates rates of streamflow for all riverbed nodes in an iterative, upstream-to-downstream method. Starting with input streamflow at the uppermost stream node and working downstream, the flow is calculated for each river reach on the basis of incoming flow (from the upstream reach) and the gain from or loss to the aquifer throughout the length of the reach. In addition, the model has been modified to allow for incoming streamflow from tributaries along the river. The model computes river stage for each reach using the stage-discharge relationship established for each reach by the user. The stream-aquifer gains and losses then are calculated as a function of streambed area and leakance and the gradient between the stream and the aquifer. The allowable head difference between the aquifer and the stream is limited when the aquifer head is below the base of the stream, as indicated in equation 2b. The algorithm for this is located in subroutine river, lines QRV 1010-1330.

The model uses a budget algorithm to account for water applied at the land surface (incident water), such as water from precipitation or irrigation. Incident water is accounted for in the model either as deep percolation (aquifer exchange) or as evapotranspiration. Direct runoff is not accounted for in the present program. The relative amounts of incident water that are simulated as eventually going to either deep percolation or evapotranspiration depend on the rate of incident water application, the rate of consumptive-use demand, and the moisture-holding capacity of the soil--all factors entered by the user. Simulated precipitation is distributed equally to all elements of the finite-element grid, and irrigation water originating from well pumpage is directed to elements as specified by the user.

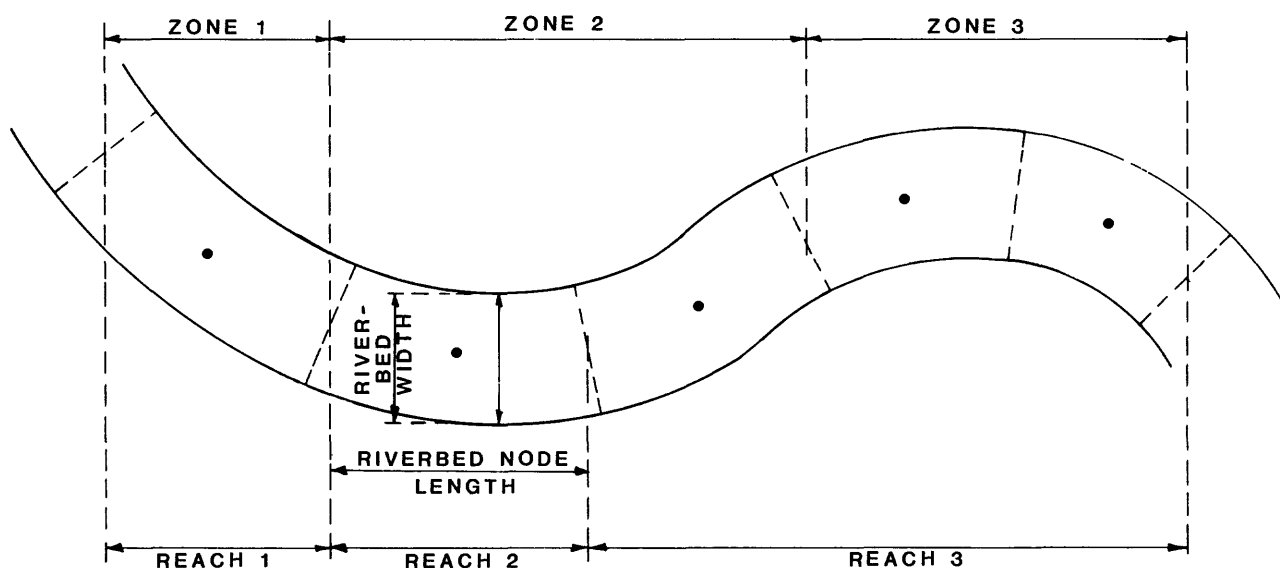


Figure 21.--Method of describing riverbed characteristics for model input.

The model uses a single composite consumptive-use demand for all model nodes for any given month of simulation. This value is determined outside the model utilizing appropriate methods, such as the Blaney-Criddle formula. The model computes the amount of water available to meet the consumptive water requirements for crops using an algorithm developed by the U.S. Department of Agriculture (1967, p. 27) and then determines evapotranspiration. In model areas where the ground-water levels are within a depth criteria set by the user, additional "surface-induced evapotranspiration" is calculated.

Output from the model consists of a water budget, constant-head and stream-aquifer fluxes, water levels and drawdowns from the starting water levels, streamflow by node, and, if selected, a plot file. A mass balance is computed for the model from all sources with the results expressed in two ways: (1) as rates for the current time step and (2) as a cumulative volume of water for each source from the start of the simulation. A mass residual is determined from the differences between the sum of sources and the sum of discharges. Printouts of constant-head and stream-aquifer fluxes by node have been added to the model. The plot-file output is in standard Calcomp¹ format, except that the plot initiation statement has been changed to comply with the format required for a Zeta plotter.

Sample Problems

The 20 x 20 node grid shown in figure 20 of this attachment was utilized in the following sample problems (except for problem 5) with model setup changed for each problem to accommodate boundaries, rivers, and wells. Grid spacing is 1,000 feet.

Problem 1: Determine radial flow to a well in a confined aquifer for the case of an infinite aquifer with no leakage. A pumping well is located next to the node on row 11, column 10 (fig. 22). Aquifer properties are held constant, and no recharge or boundary-flow conditions are applied. The normal model boundaries are no-flow boundaries, unless altered with other model options. However, the aquifer is assumed to be infinite; therefore, the model boundaries were set at a distance to eliminate any significant effects on the results of the problem. An analytical solution originally by Theis (Lohman, 1972, p. 15) is used for comparison.

$$s = \frac{Q}{4\pi T} W(u) \quad (3)$$

¹ The use of brand names in this report is for identification purposes only and does not imply endorsement by either the U.S. Geological Survey or the Kansas State Board of Agriculture.

where $u = \frac{r^2 S}{4Tt}$;
 Q = constant discharge rate from well [L^3/t];
 T = transmissivity [L^2/t];
 r = distance from discharging well to point of observation [L];
 S = storage coefficient [dimensionless];
 t = time since pumping began [t];
 s = drawdown [L]; and
 $W(u) = \int_u^\infty \left(\frac{e^{-u}}{u}\right) du$.

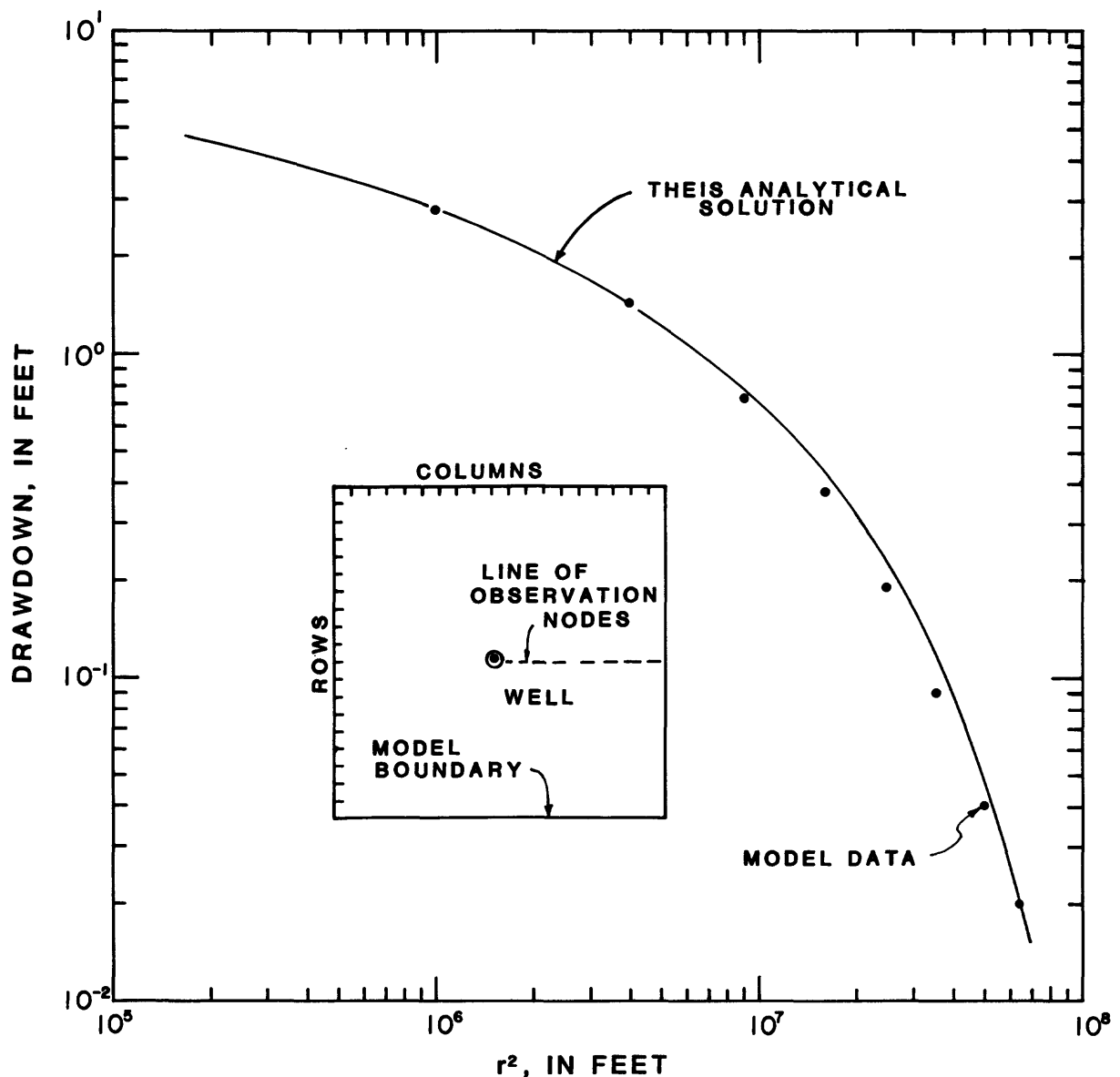


Figure 22.--Numerical simulation of infinite aquifer with Theis analytical-solution assumptions.

Aquifer and model properties are as follows:

Initial potentiometric surface	100 feet
Storage coefficient	0.01
Transmissivity	10,000 ft ² /d
Well pumping rate	1.3369 x 10 ⁵ ft ³ /d
Time step	1 day
Time of solution	6 days

Problem 2: Determine radial flow to a well in a confined aquifer for the case of an infinite aquifer with constant-head boundary and no leakage. A pumping well is located next to the node on row 2, column 10 (fig. 23) and is 2,000 feet from the constant-head boundary. The observation node is located on row 7, column 10. Aquifer properties are held constant. The aquifer is assumed to be infinite, except for the constant-head boundary, and the model boundaries were set at a distance to eliminate any significant effects on the results of the problem. Analytical solutions of the image-well theory (Lohman, 1972, p. 60) and of Theis for the case of an infinite aquifer, as described before, are used for comparison. As described by Lohman (1972), a solution of the single-boundary problem utilizing a discharging or recharging image well can be obtained as follows:

$$s_o = \frac{Q}{4\pi T} [W(u) \pm W(u_i)] \quad (4)$$

where

s_o = algebraic sum of the drawdown of the pumping and image well;

$$u_p = \frac{r_p^2 S}{4Tt} ;$$

$$u_i = \frac{r_i^2 S}{4Tt} ;$$

r_p = the distance from the pumped well to the observation well [L];

r_i = the distance from the observation well to the image well [L]; and

other variables are as defined before.

Aquifer model properties are as follows:

Initial potentiometric surface	200 feet
Storage coefficient	0.01
Transmissivity	5,000 ft ² /d
Well pumping rate	1.3369 x 10 ⁵ ft ³ /d
Time step	1 day
r_p	4,000 feet
r_i	8,000 feet

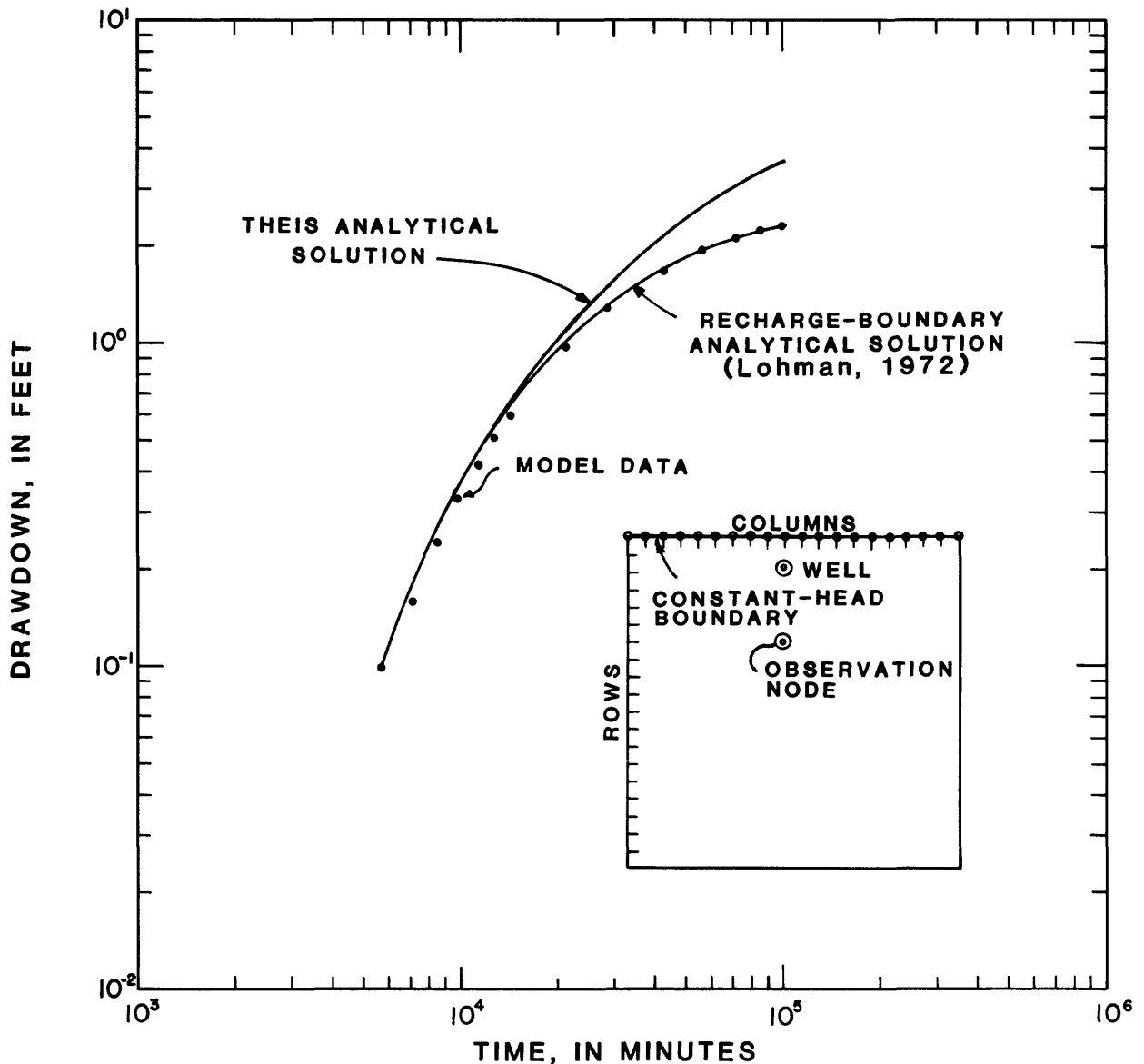


Figure 23.--Numerical simulation of a discharging well in an aquifer bounded with a constant-head boundary.

Problem 3: Determine radial flow to a well in a confined aquifer with an impermeable boundary and no leakage. A pumping well is located next to node on row 2, column 10 (fig. 24) and is located 2,000 feet from the boundary. The observation node is located on row 7, column 10. Aquifer properties are held constant. The upper boundary is a no-flow boundary. The remainder of the aquifer is assumed to be infinite in extent, and the other model boundaries are set at a distance to eliminate any significant effects on the results of the problem. Analytical solutions of Theis and of the image-well theory, as described previously, were used for comparison.

Aquifer and model properties are as follows:

Initial potentiometric surface	200 feet
Storage coefficient	0.01
Transmissivity	5,000 ft ² /d
Well pumping rate	1.3369×10^5 ft ³ /d
Time step	1 day
r_p	4,000 feet
r_i	8,000 feet

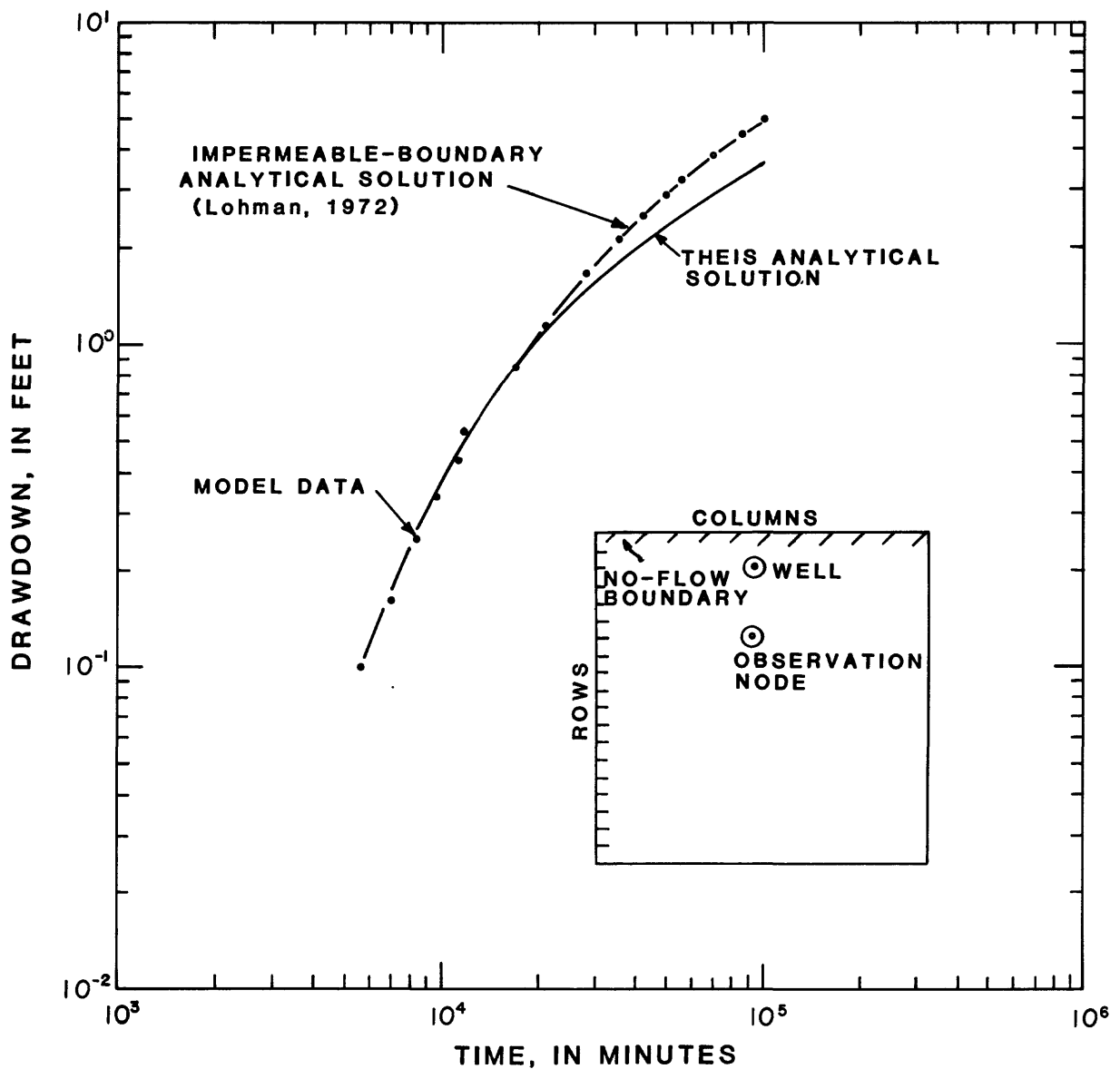


Figure 24.--Numerical simulation of a discharging well in an aquifer bounded with an impermeable boundary.

Problem 4: Determine radial flow to a well in a confined aquifer with induced infiltration from a stream (fig. 25). Pumping well is located next to a node on row 13, column 9. Stream is located along row 11. The observation node is located on row 17, column 9. Flow within the stream was set at a value much greater than the flow loss to the aquifer, and the stage-discharge rating was set so that any flow in the stream would maintain the same stage in the river.

In this case, a comparison of model results is made for a problem in which there is no exact analytical solution. Although an analytical solution to this problem is not available, it is believed that the exact unknown solution should be bounded by the two cases given below until the computed head changes are effected by the model no-flow boundaries. Model boundary conditions are not significant during the time simulated for the example.

Analytical solutions of Theis for the case of an infinite aquifer and of the image-well theory, as described previously, were used for comparison of results.

Aquifer and model properties are as follows:

Initial potentiometric surface	200 feet
Stream stage	200 feet
Storage coefficient	0.01
Transmissivity	10,000 ft ² /d
Well pumping rate	2.6738×10^5 ft ³ /d
Time step	1 day
r_p	4,000 feet
r_i	8,000 feet
K_i	1 ft/d
b_i	1 foot
Riverbed elevation	199 feet
Stream width	100 feet

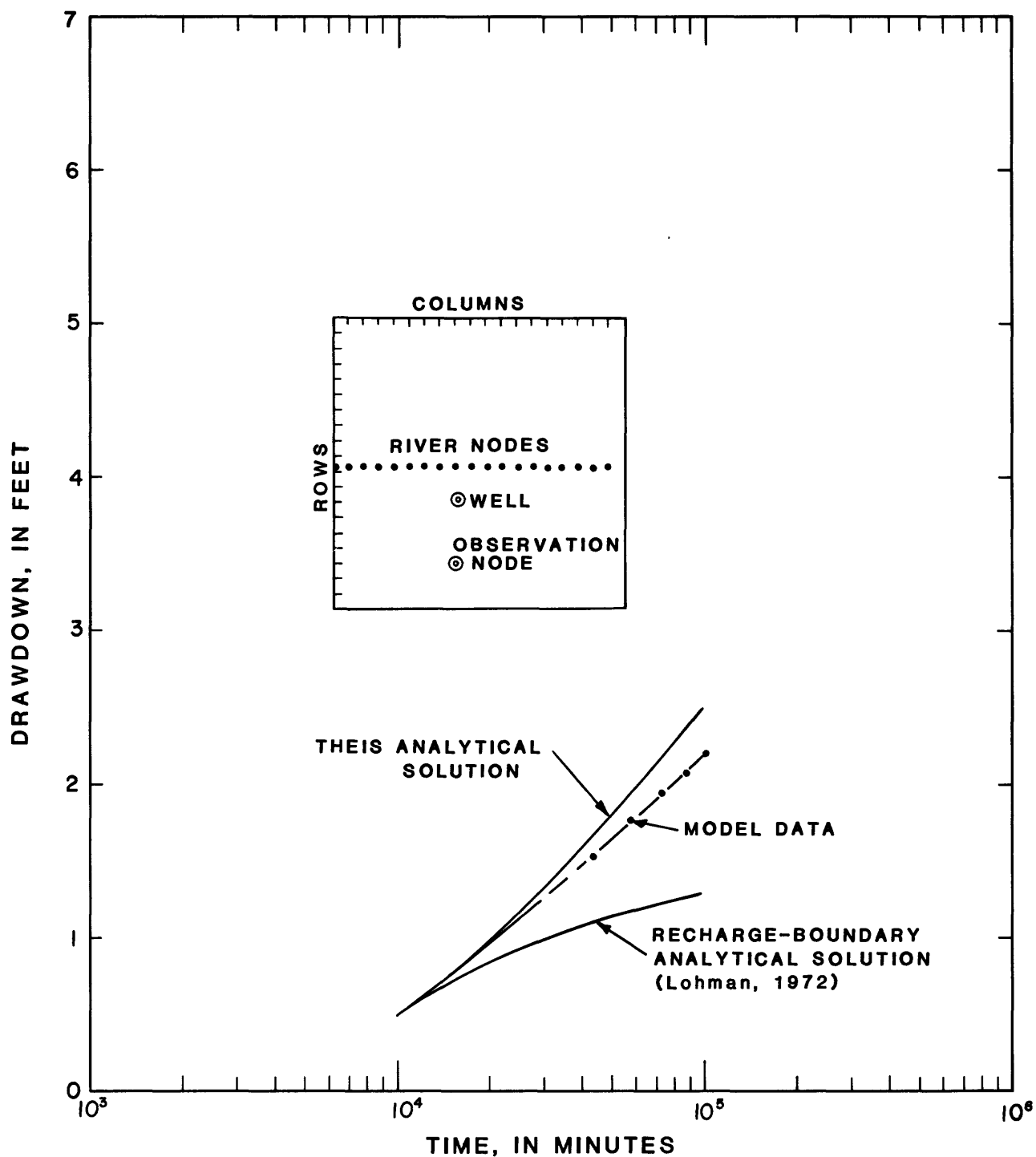


Figure 25.--Numerical simulation of a discharging well with induced infiltration.

Problem 5: Water-table aquifer with constant-head boundaries and constant recharge--steady-state conditions (fig. 26). Streams (constant-head boundaries in model) are located 16,000 feet apart. Aquifer is assumed to be homogeneous and isotropic.

Analytical solution for an aquifer under steady-state conditions, uniform recharge, and bounded on two sides by fully penetrating parallel streams of infinite length (fig. 27), as shown in Ferris and others (1962, p. 131), was used for comparison of results. The steady-state, ground-water-level profile is defined as:

$$h_o = \left(\frac{a^2 W}{2T} \right) \left(\frac{2x}{a} - \frac{x^2}{a^2} \right) , \quad (5)$$

where

h_o = elevation of the water table at observation point x , with respect to the mean stream level [L];

W = constant rate of recharge to the water table [L/t];

a = distance from the stream to the ground-water divide [L]; and

x = distance from the stream to the observation point [L].
The stream on the left is assumed to be the reference point for x .

Aquifer and model properties are as follows:

Initial potentiometric surface	40 feet
Stream stage	40 feet
Transmissivity	34,560 ft ² /d
Recharge	8.06 x 10 ⁻³ ft/d
a	8,000 feet

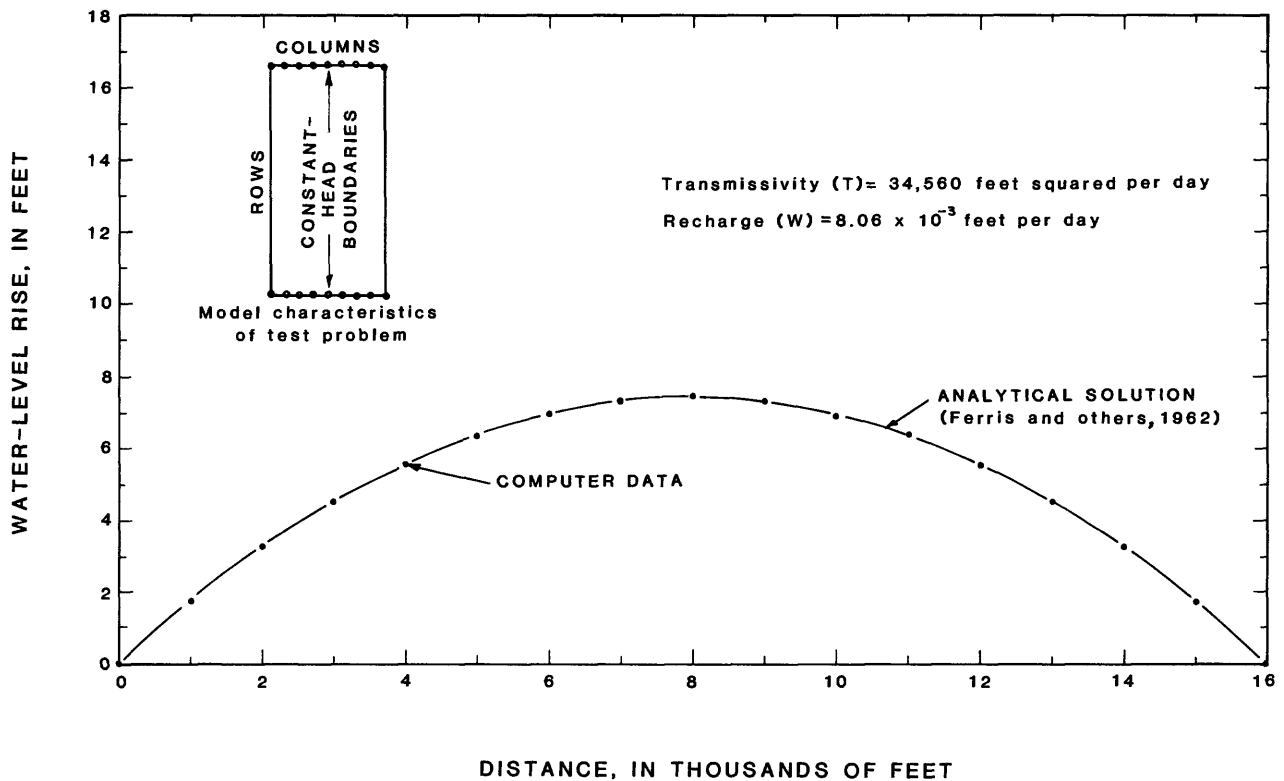
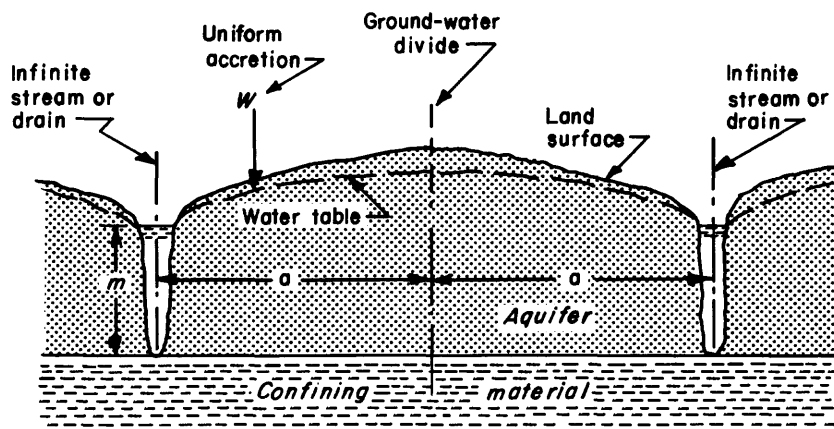


Figure 26.--Numerical simulation of water-table aquifer with constant-head boundaries and constant recharge.

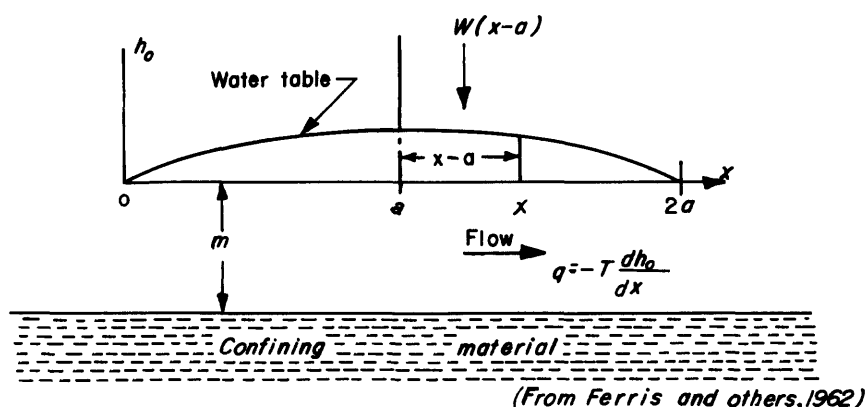
Description of Program Routines

The finite-element areal-flow model consists of the main program and 25 subroutines. The Tracy model was modified from 16 subroutines with 10 entry points to 25 subroutines for use on the Prime computer. They are listed in the order that they are first called. The function of each routine is outlined following this list.

- | | | |
|-----------------|------------|----------------|
| 1. MAIN PROGRAM | 10. LEMENQ | 19. SSOLVE |
| 2. DATAIN | 11. LEMENT | 20. ADCOMP |
| 3. DATE | 12. QRIVER | 21. ASOLVE |
| 4. NWDATE | 13. RIVER | 22. QZEE, |
| 5. ARRAY | 14. PTFNDR | 23. PUMPNU |
| 6. LAMBRT | 15. BNDCON | 24. LOADQR |
| 7. PLTG | 16. QBOUND | 25. OUTPUT |
| 8. PLTR | 17. MASBAL | 26. BLOCK DATA |
| 9. PLTW | 18. SDCOMP | |



A. SECTION VIEW OF IDEAL AQUIFER SUBJECT TO UNIFORM ACCRETION, BOUNDED BY PARALLEL STREAMS



B. NOMENCLATURE FOR MATHEMATICAL ANALYSIS OF PROFILE SHOWN IN SKETCH A

Figure 27.--Section views for analyzing steady-state flow in hypothetical aquifer of large thickness with uniform accretion from precipitation.

Main Program

This part of the model establishes the size of the arrays needed for the problem. It controls the logical flow of the program based on the options selected by the user and serves as the driver for all subsequent subroutines. The main program contains two principal loops. The first controls the progression through time in the transient-flow simulation. The second controls the iterative-solution technique for modeling stream-aquifer interaction.

Subroutine DATAIN

This subroutine handles all input and output related to the problem specification and aquifer properties. The required input data and format are detailed in the next section. The routine also generates the finite-element grid and calls the grid-plotting routines.

Called by: MAIN PROGRAM

Subroutines called: DATE, ARRAY, LAMBRT, PLTG

Subroutine DATE

The call to this routine initializes the data needed to keep track of calendar dates for the transient simulation.

Called by: DATAIN

Subroutine NWDATE

This routine updates the calendar dates at the end of each time step.

Called by: MAIN PROGRAM

Subroutine ARRAY

This routine reads in the arrays of aquifer properties that are defined at each node. It allows for flexibility as to where the data will be read from and the format required. More details can be found in the next section.

Called by: DATAIN

Subroutine LAMBRT

This routine converts all input data read in as latitude and longitude into x- and y-coordinate values. The algorithm uses a polynomial approximation to the Lambert conical projection with an origin of 39° latitude and 120° longitude.

Called by: DATAIN, QRIVER, PUMPNU

Subroutine PLTG

This subroutine uses standard CALCOMP calls to plot the finite-element grid based on the input data. Use of this routine is optional, but it is the easiest means of verifying the accuracy of the input.

Called by: DATAIN

Subroutine PLTR

This routine plots the river nodes on the finite-element grid.

Called by: QRIVER

Subroutine PLTW

This routine plots the wells on the grid.

Called by: PUMPNU

Subroutine LEMENQ

This routine performs the integration of the flow equation over a linear, quadrilateral finite element. The integrations for both the steady and transient terms are done using a four-point, Gaussian quadrature technique, and these terms then are loaded into the matrix A.

Called by: MAIN PROGRAM

Subroutine LEMENT

This routine performs the integration of the flow equation over a linear triangular element. The integrations are done analytically, and both the steady and transient terms are loaded into the matrix A.

Called by: MAIN PROGRAM

Subroutine QRIVER

This routine reads in all the data to be used in calculating stream-aquifer interaction. It also loads the leakance terms for the stream into the matrix A.

Called by: MAIN PROGRAM

Subroutine RIVER

This routine routes the daily inflow to the stream through the channel and determines the amount of leakage into or out of the stream based on the best values of aquifer head. As values of aquifer head at the next iteration level are determined with the river leakage handled implicitly, the routine is called again. This procedure is repeated until convergence is achieved.

Called by: MAIN PROGRAM

Subroutine PTFNDR

This routine determines the local coordinates and appropriate integrals of the basis function for a specified point within an element, given the global coordinates of the point and of the nodes defining the element. For triangular elements, the coordinates can be solved directly in terms of the basis functions. For the quadrilateral elements, the two quadratic equations of global position must be solved simultaneously to yield the local coordinates.

Called by: QRIVER, PUMPNU

Subroutine BNDCON

This routine saves the matrix equation, formulated in matrix A, for nodes having a specified head. It then applies the boundary condition to these nodes by zeroing out the row and column of the matrix and placing a 1.0 on the diagonal.

Called by: MAIN PROGRAM

Subroutine QBOUND

This routine adjusts the right-hand side of the matrix flow equation to correct for the zeroing out of the symmetric terms in the rows and columns of the A matrix when applying the specified-head boundary condition. The entry also adds the specified-flux boundary terms to the R.H.S.(QRE).

Called by: MAIN PROGRAM

Subroutine MASBAL

This routine is called after the flow equation is solved. It calculates the boundary fluxes at the specified-head nodes and the storage change as a check for mass balance. These terms are printed out along with all other flow budget items that are being accumulated and the residuals of both the time-step rates (L^3/T) and the cumulative volume (L^3).

Called by: MAIN PROGRAM

Subroutine SDCOMP

This subroutine performs the decomposition of the symmetric, banded matrix A. The routine is called when the nodal transmissivities are constant and only needs to be called once.

Called by: MAIN PROGRAM

Subroutine SSOLVE

This routine does the forward and backward substitution to solve for the new head values.

Called by: MAIN PROGRAM

Subroutine ADCOMP

This routine performs the decomposition of the A matrix when it is asymmetric. This routine is called when transmissivities are head dependent and is only called once.

Called by: MAIN PROGRAM

Subroutine ASOLVE

This routine does the forward and backward substitution to solve for the new head values when the A matrix is asymmetric.

Called by: MAIN PROGRAM

Subroutine QZEE

This routine reads in climatological and soil data necessary to compute recharge due to evapotranspiration and precipitation.

Called by: MAIN PROGRAM

Subroutine PUMPNU

This routine reads in the well data, including well locations, pumpage rates, seasonal variability factors, application rates, and types and areas of application.

Called by: MAIN PROGRAM

Subroutine LOADQR

This routine loads most of the stress terms into the right-hand side of vector QRE. These terms include pumpage, recharge due to precipitation and applied water, evapotranspiration, and changes in storage due to head variation. The entry also determines how much of the pumped water gets routed back into a river.

Called by: MAIN PROGRAM

Subroutine OUTPUT

This routine prints out the new head values, cumulative drawdowns from original head values, and the river stages at specified intervals.

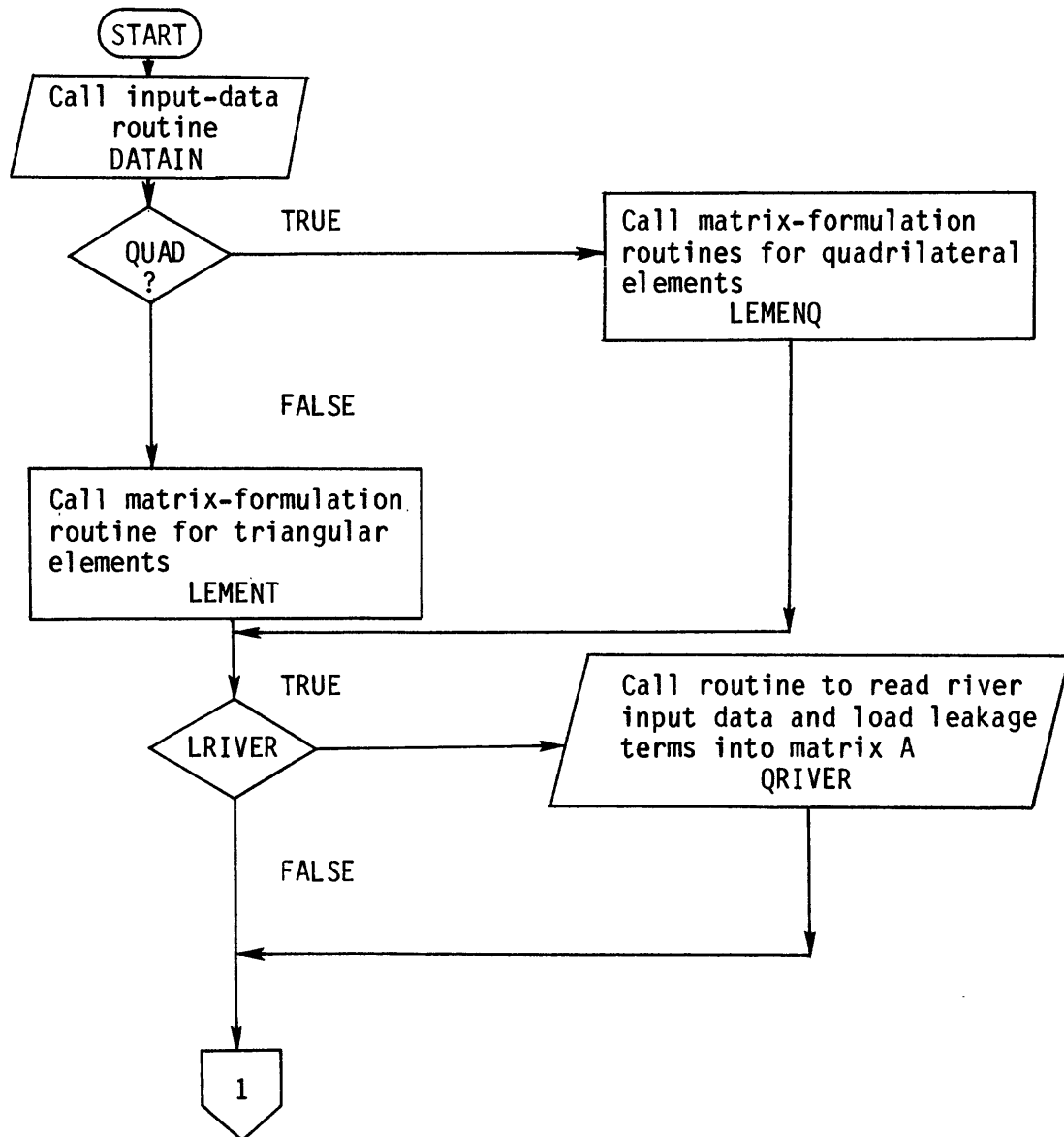
Called by: MAIN PROGRAM

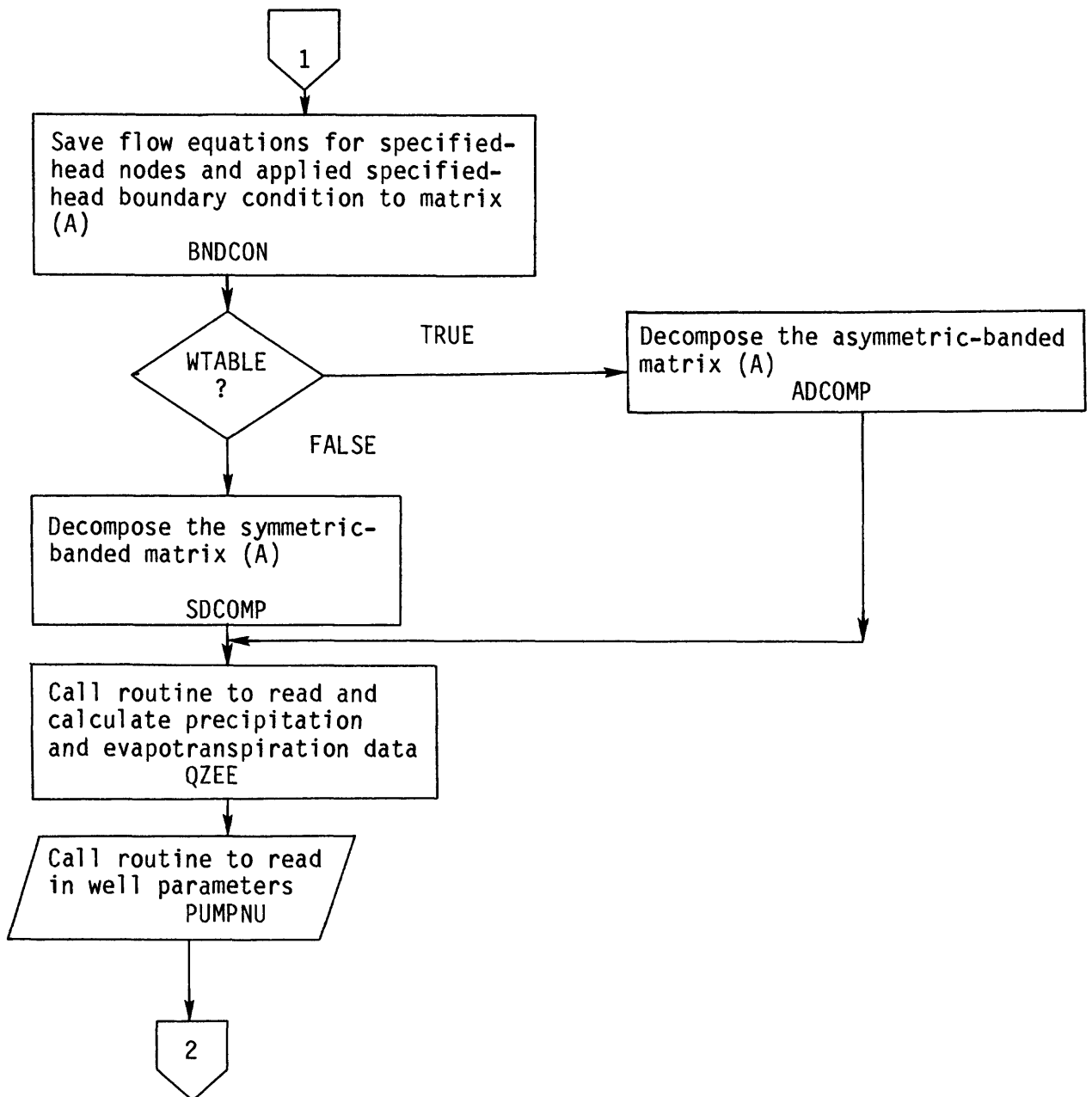
BLOCK DATA

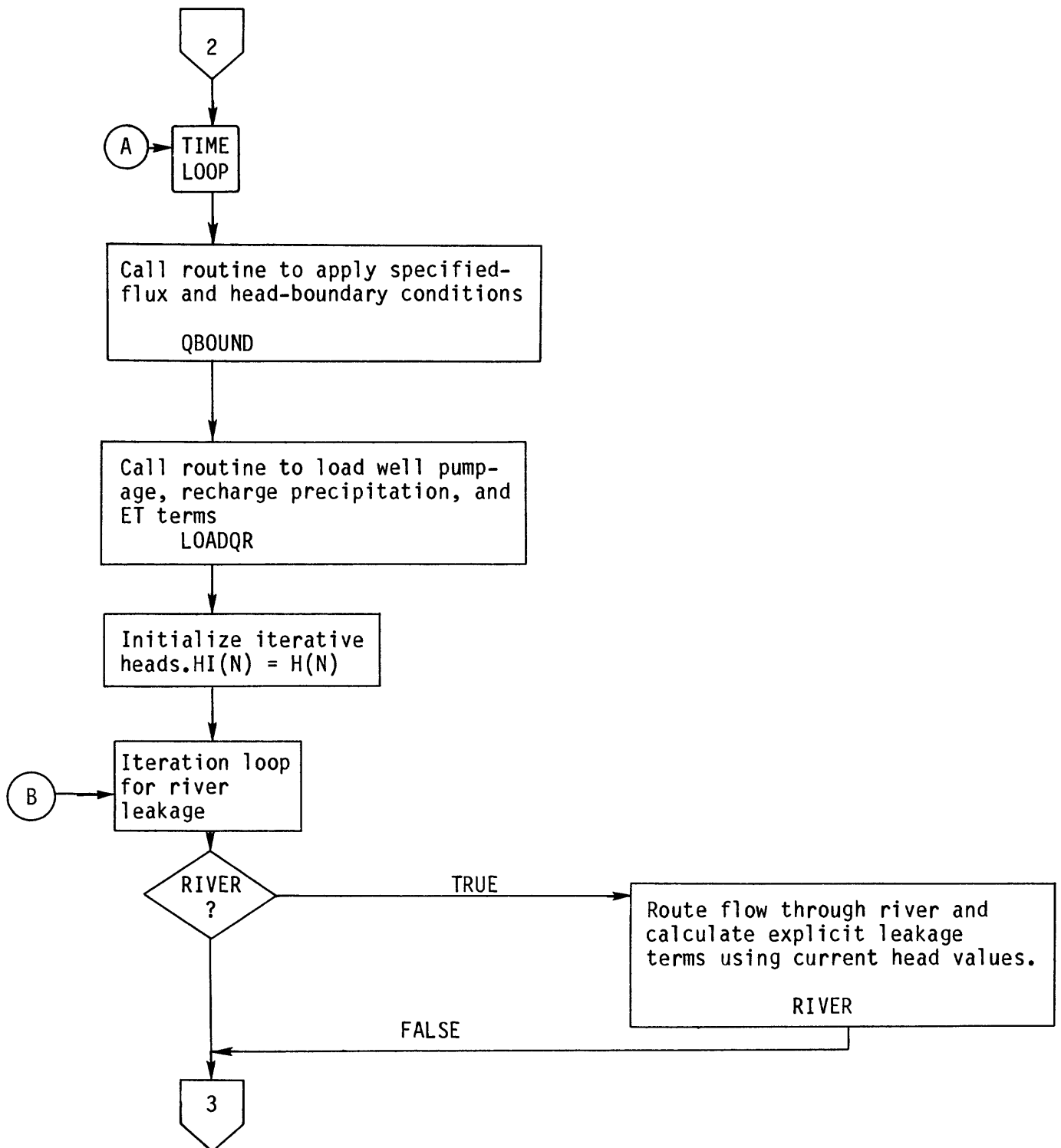
This routine is used to initialize a number of variables contained in common. These include the input and output units, evapotranspiration factors, budget-item titles, and default values for input variables that are optional.

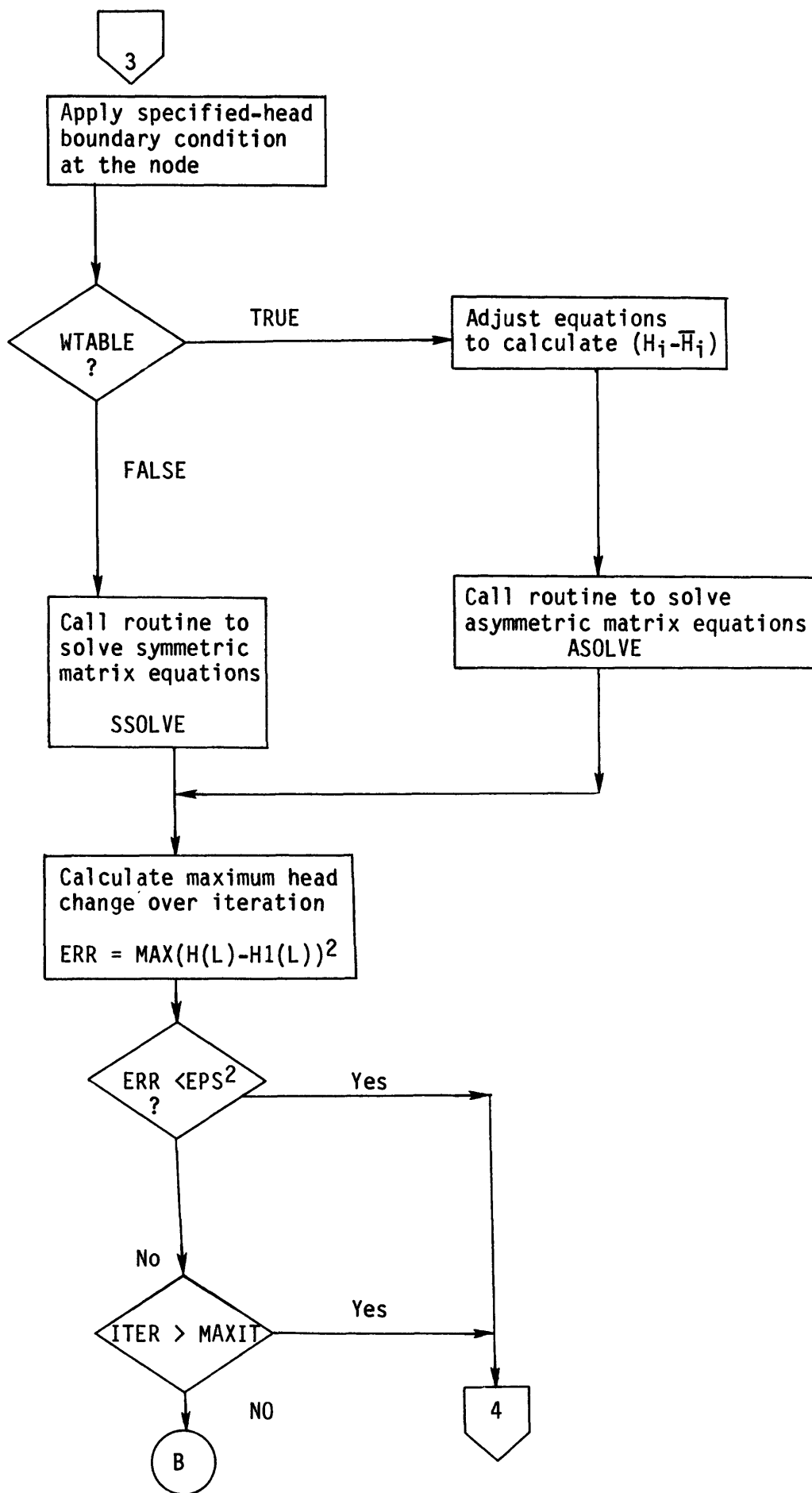
Generalized Flow Chart for Aquifer-Simulation Model

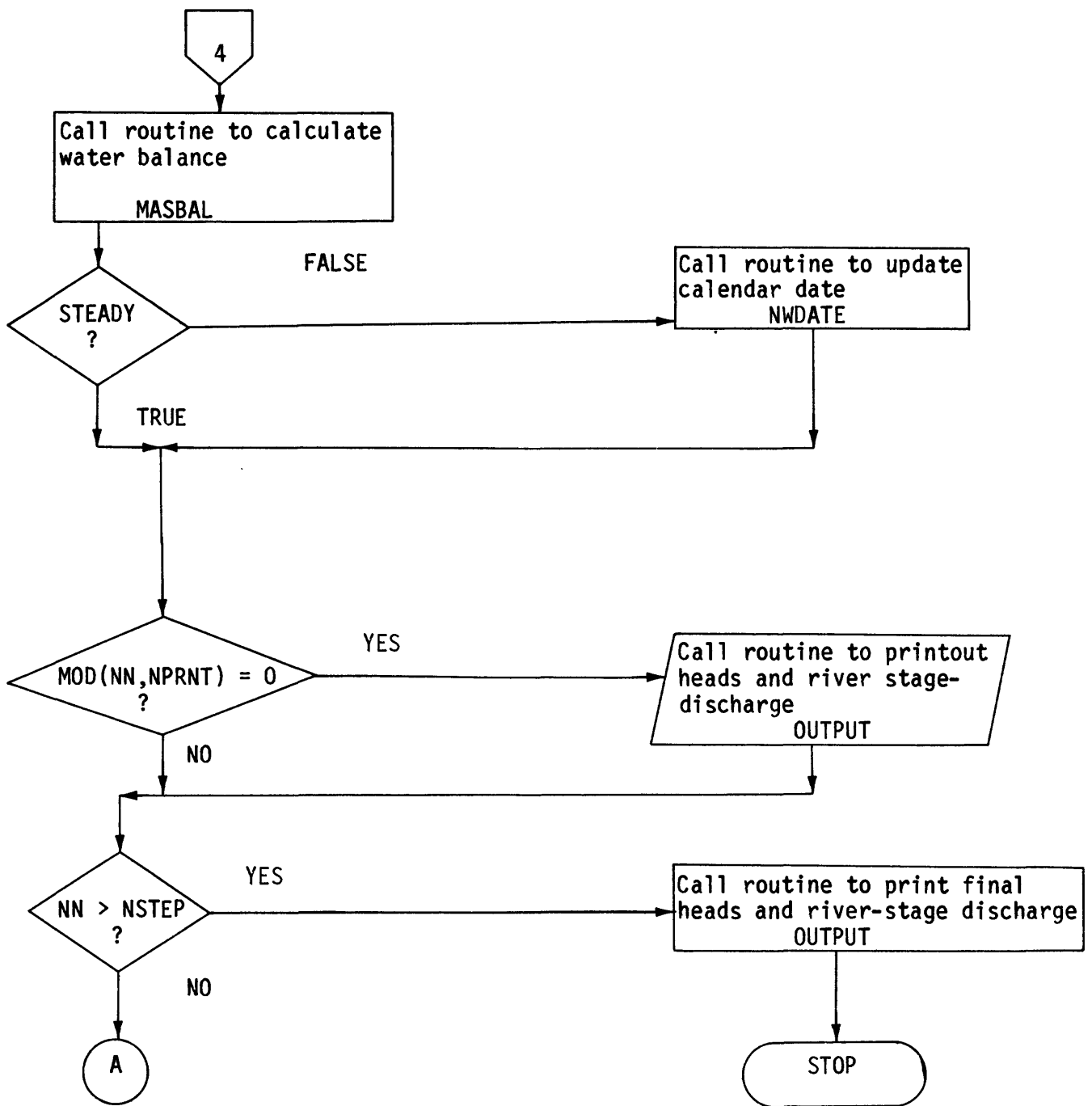
1 MAIN PROGRAM



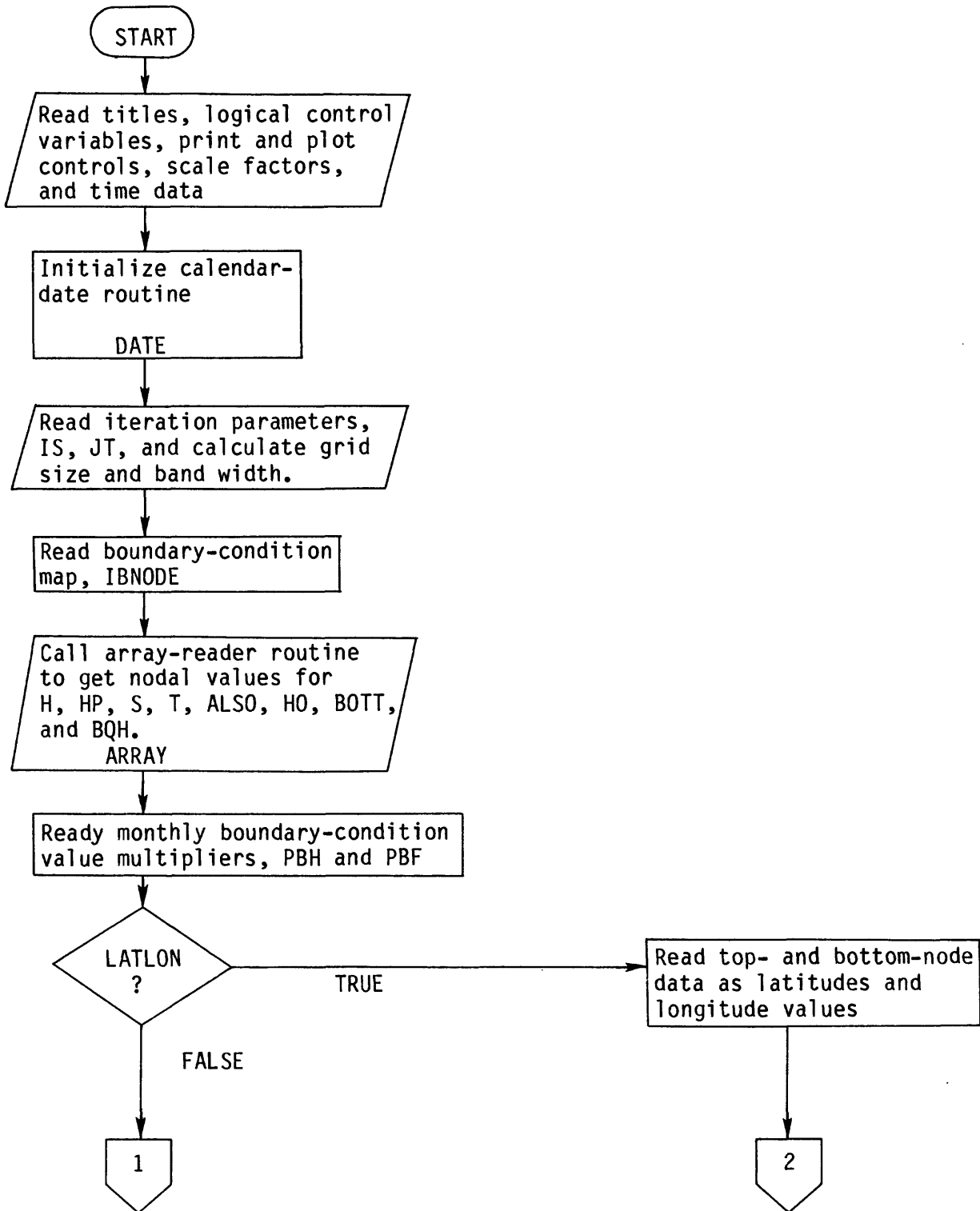


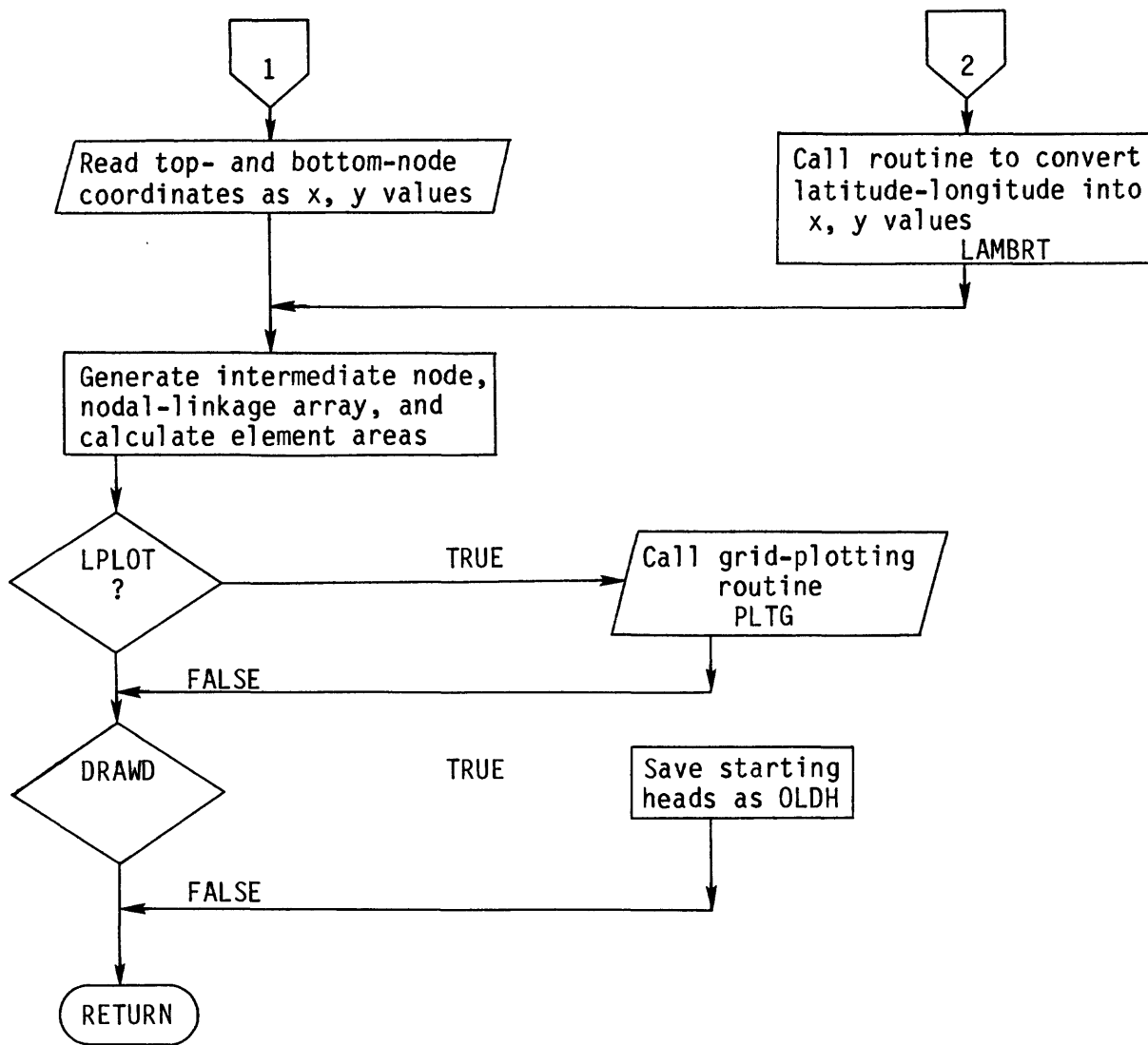




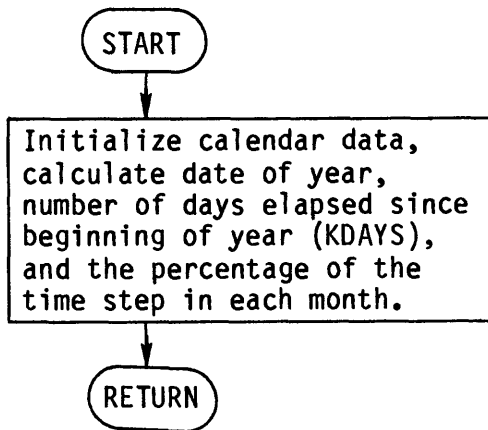


SUBROUTINE DATAIN

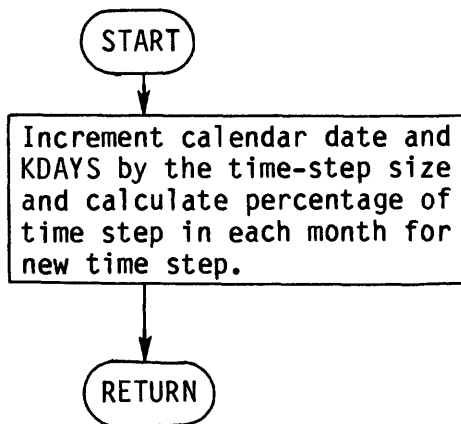




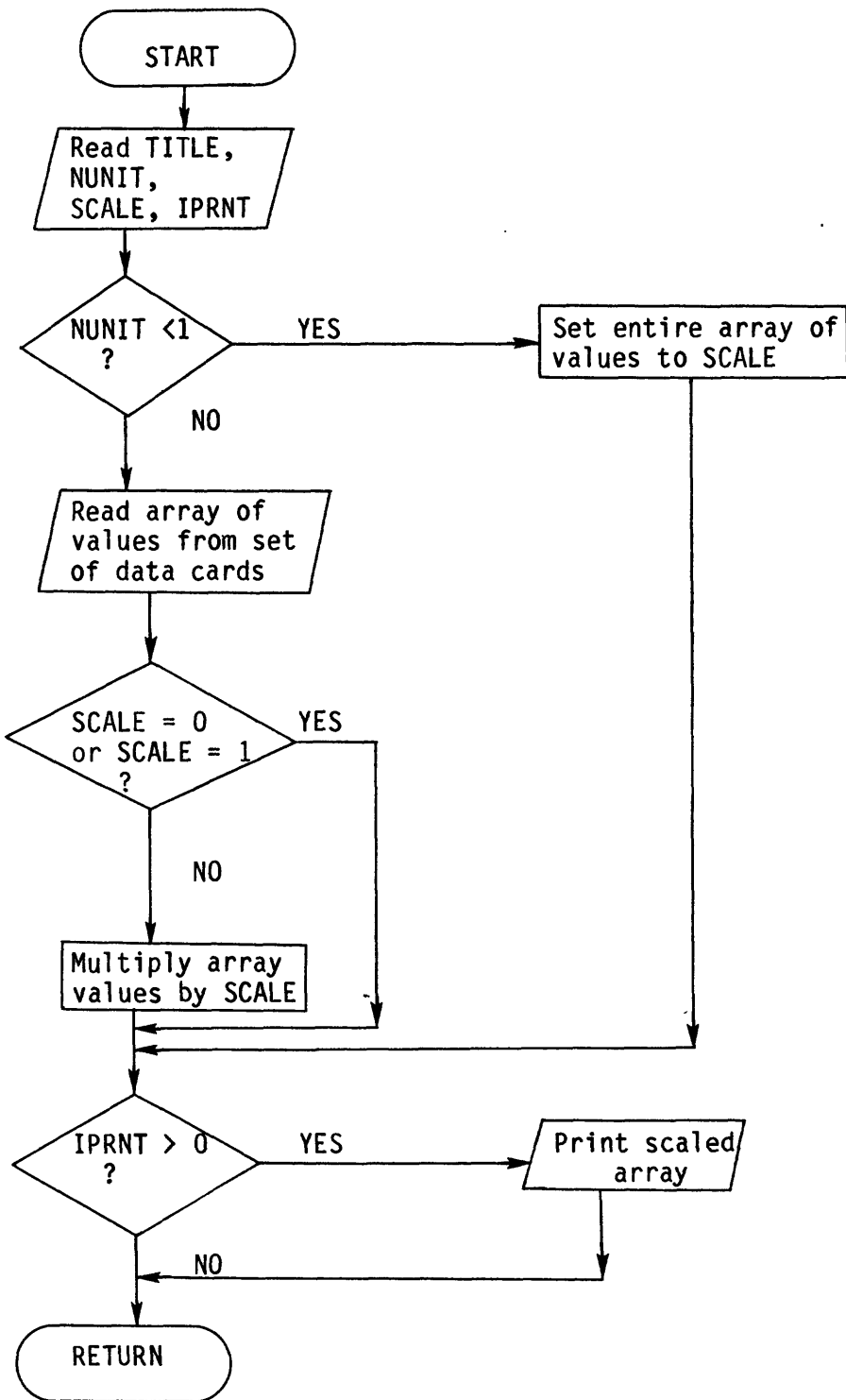
SUBROUTINE DATE



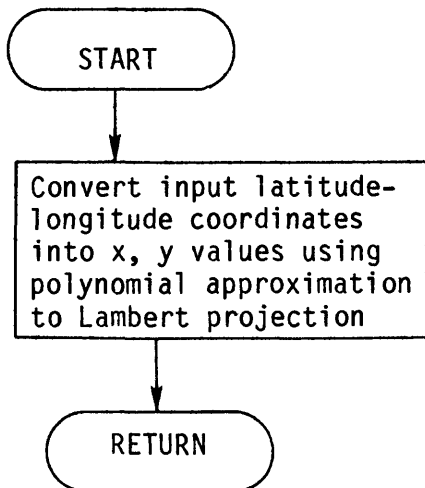
SUBROUTINE NWDATE



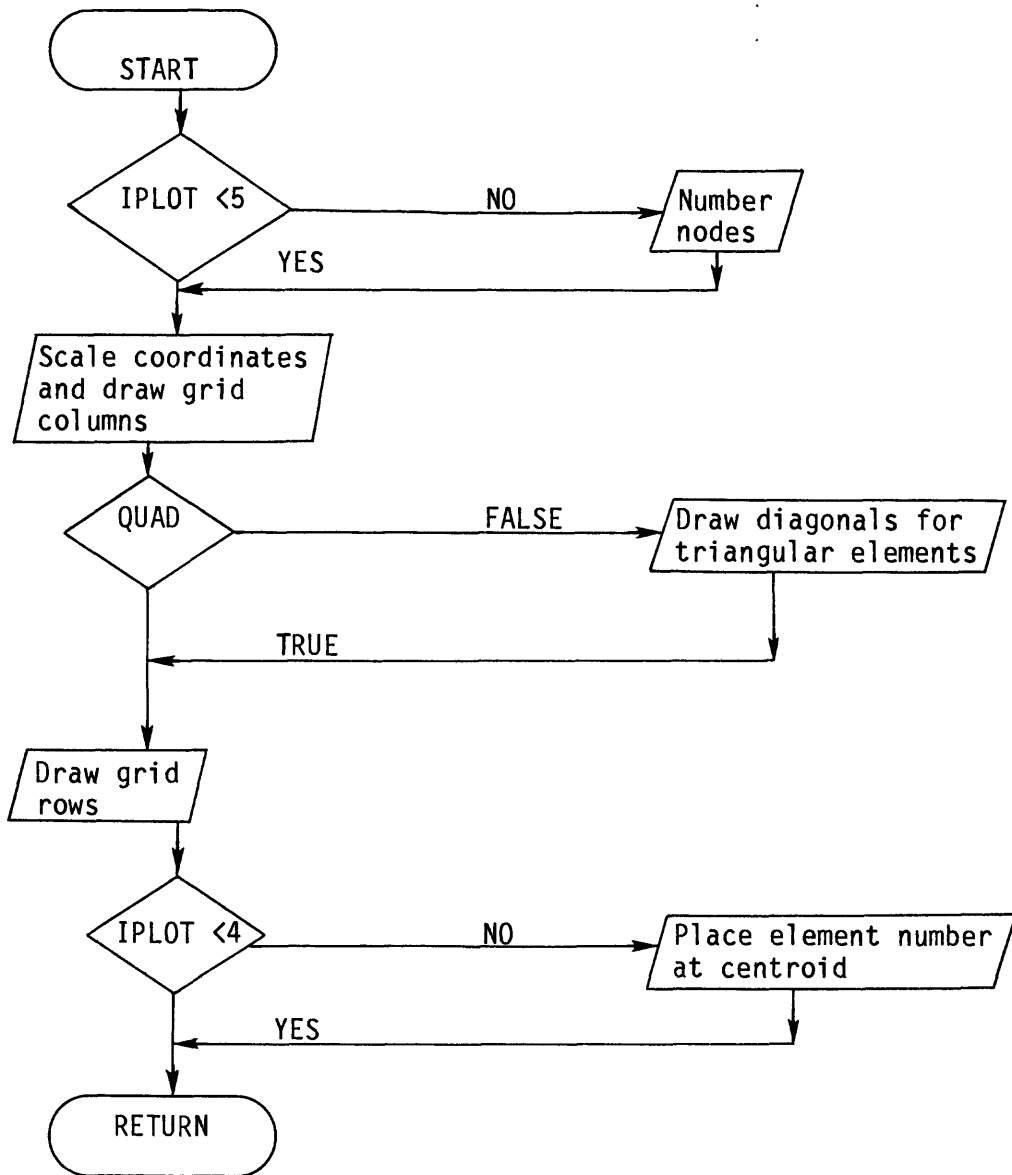
SUBROUTINE ARRAY



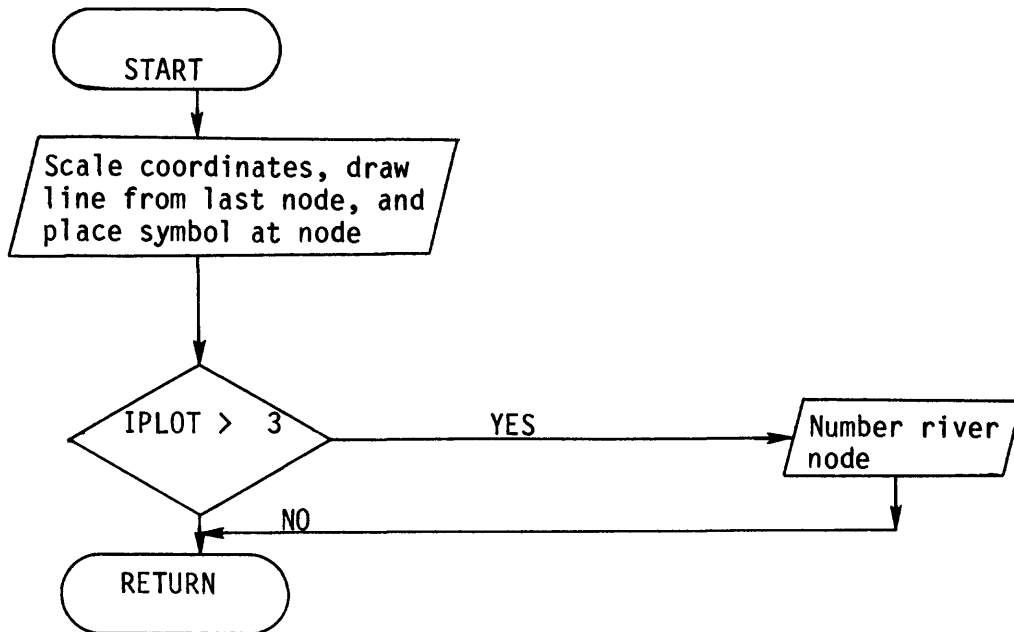
SUBROUTINE LAMBRT



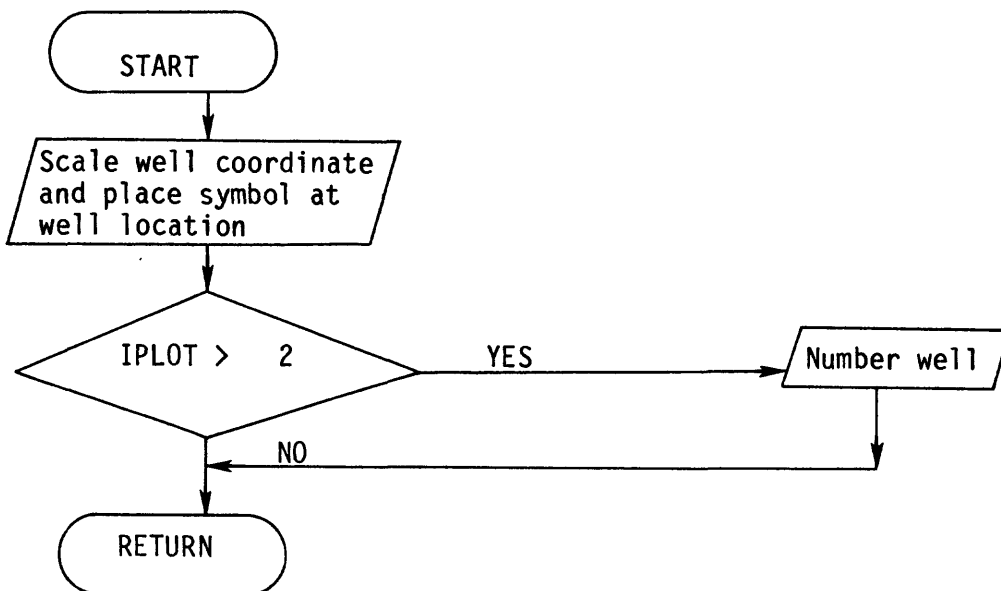
SUBROUTINE PLTG



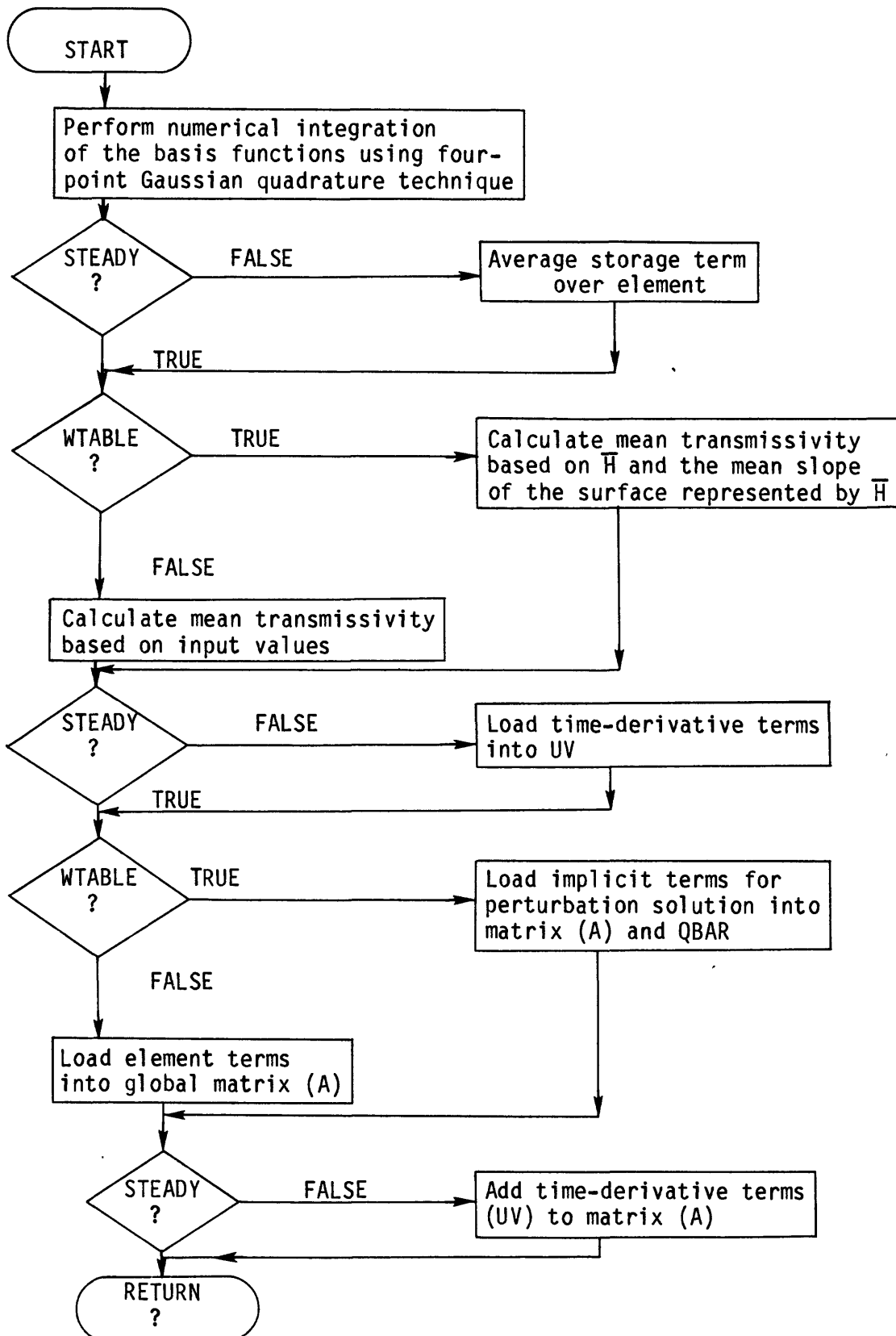
SUBROUTINE PLTR



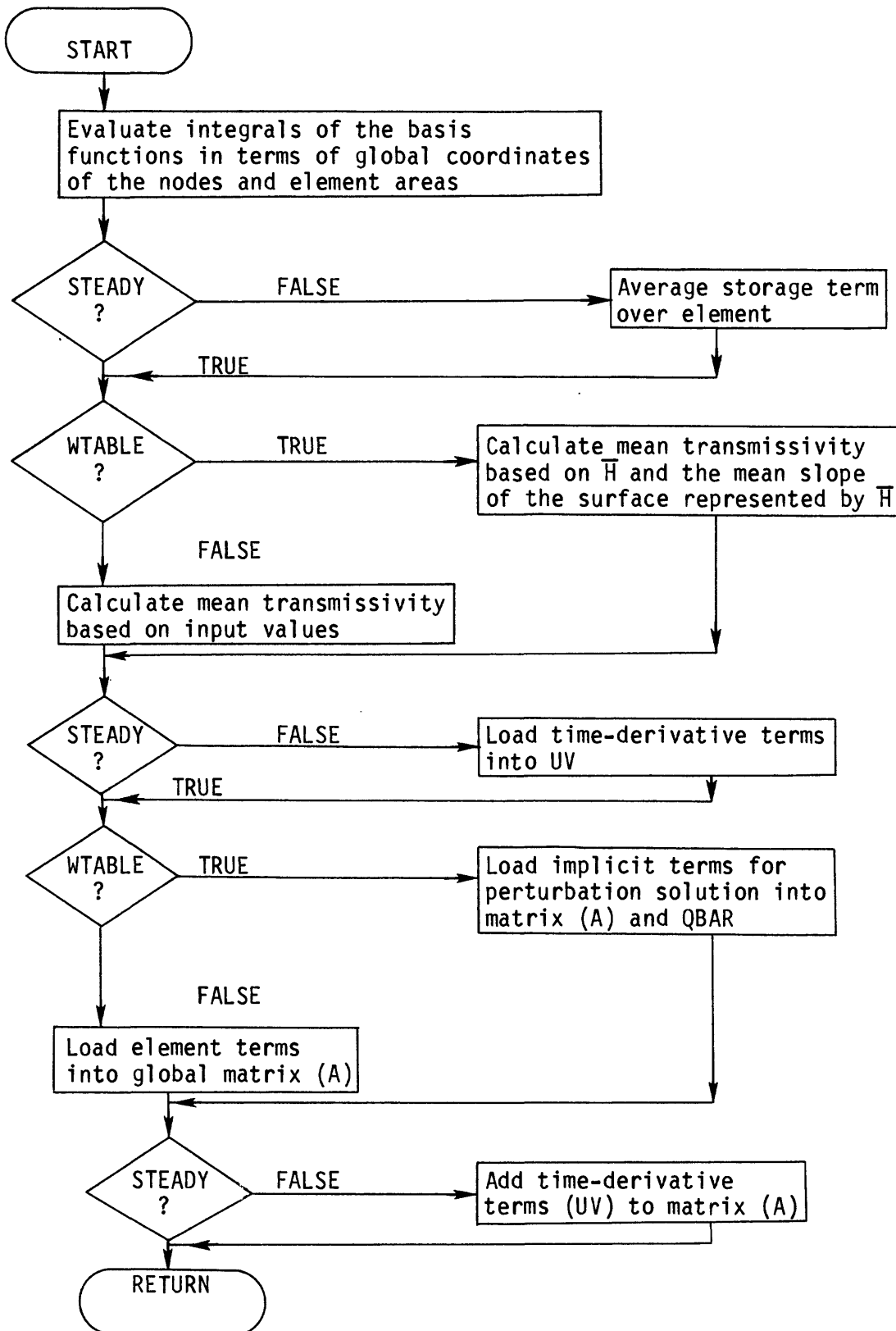
SUBROUTINE PLTW



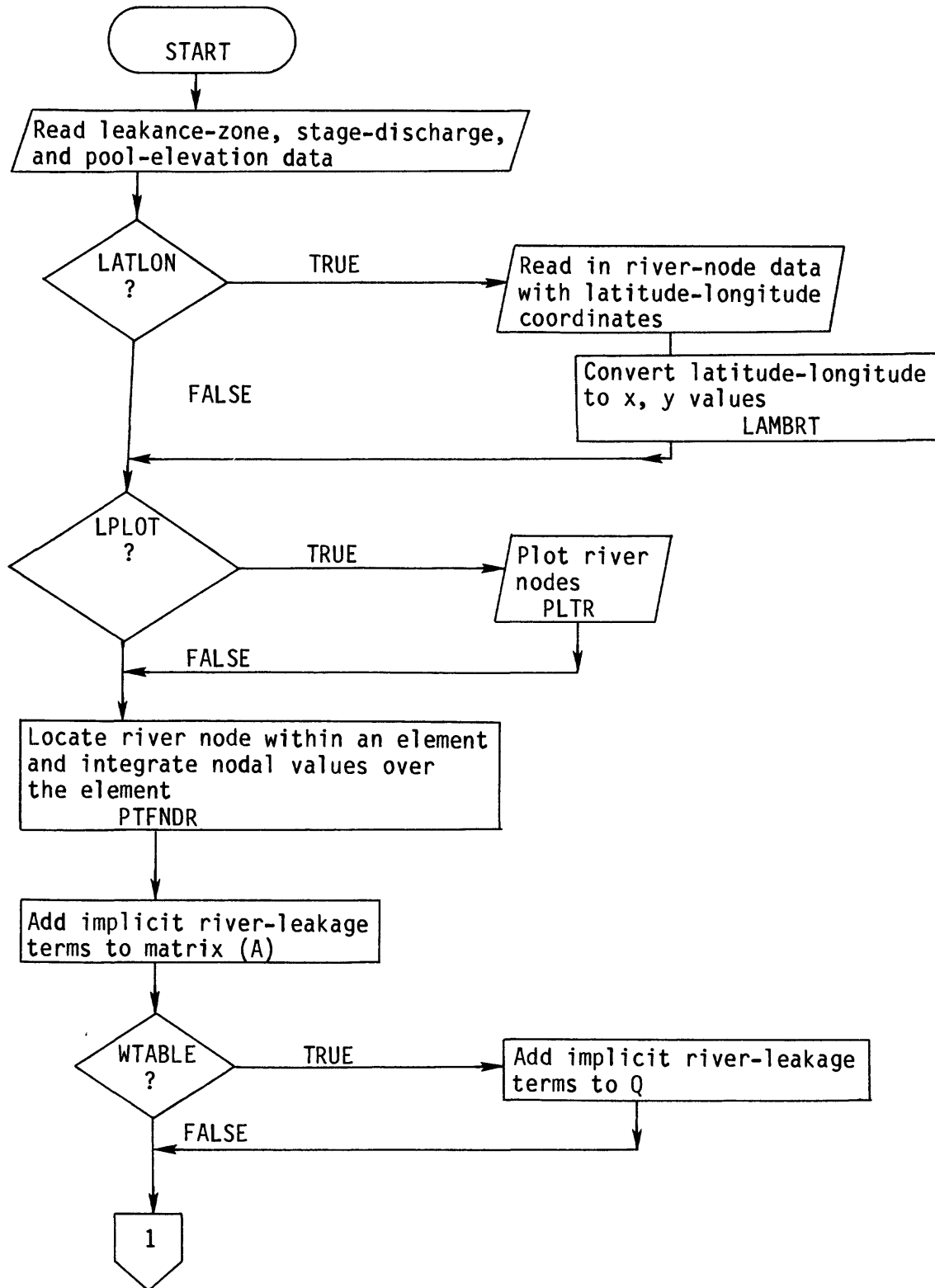
SUBROUTINE LEMENQ

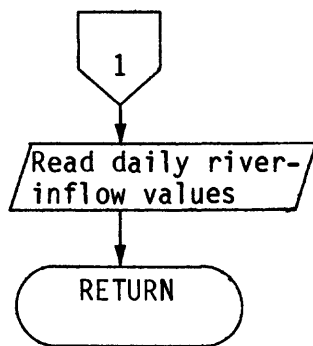


SUBROUTINE LEMENT

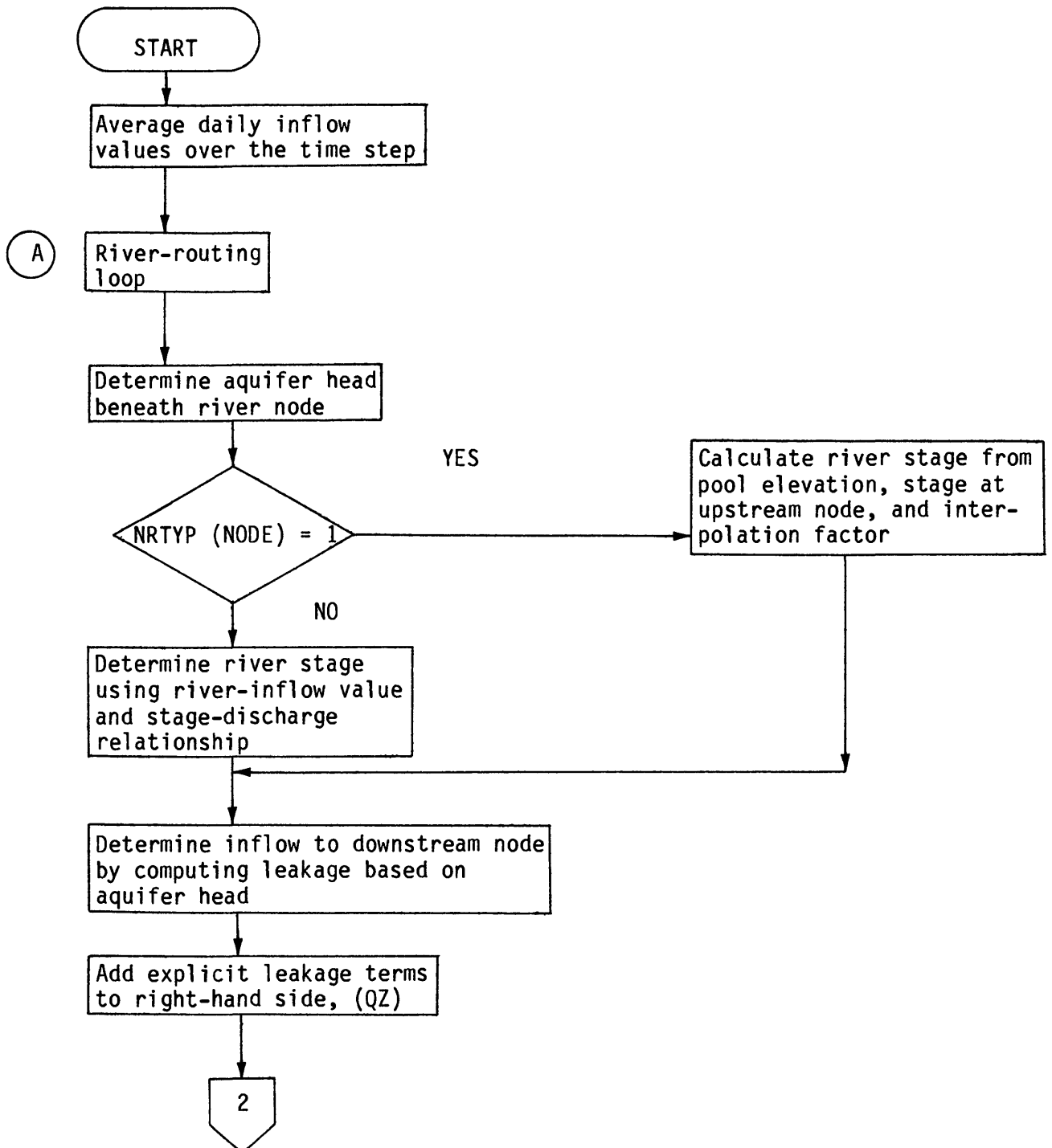


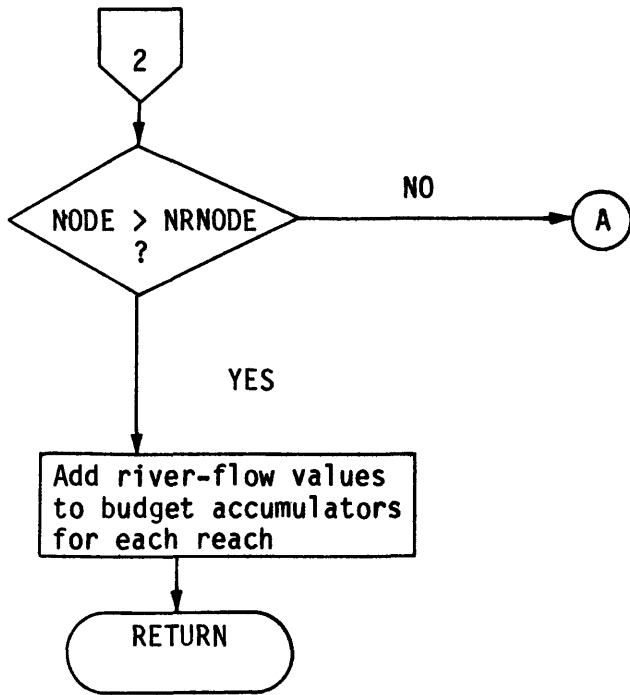
SUBROUTINE QRIVER



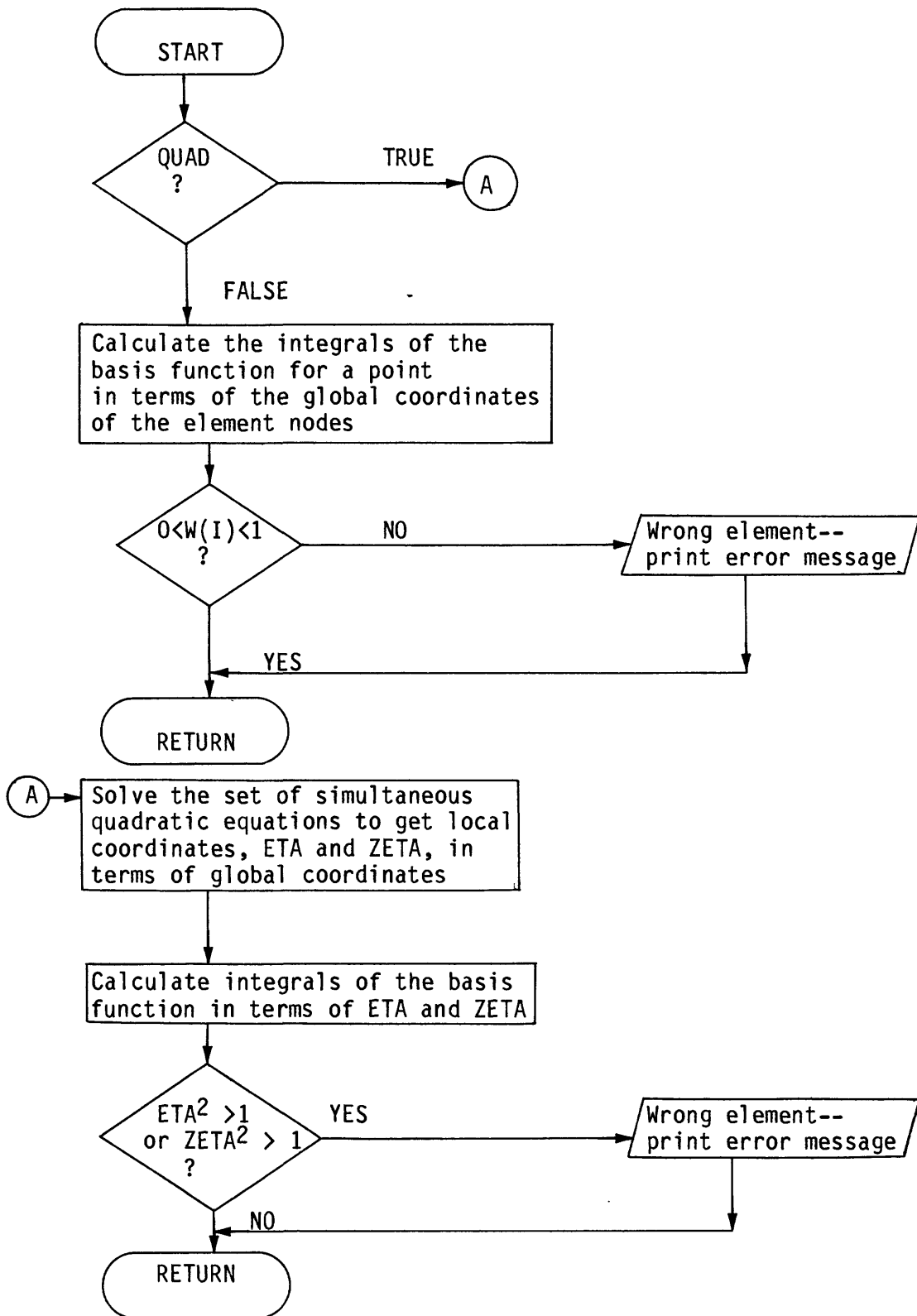


SUBROUTINE RIVER

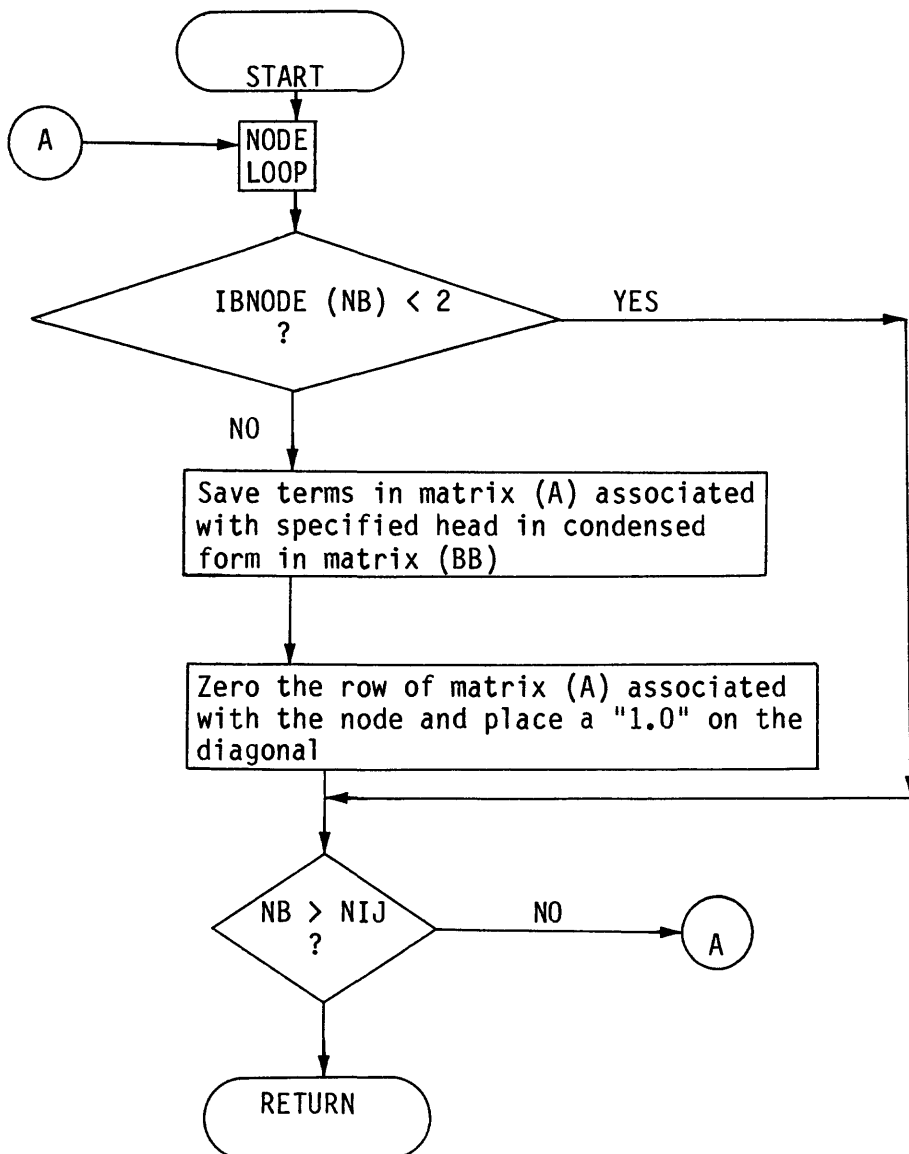




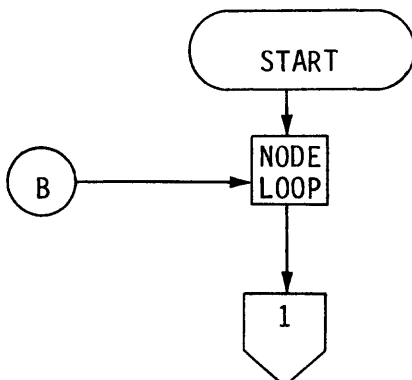
SUBROUTINE PTFNDR

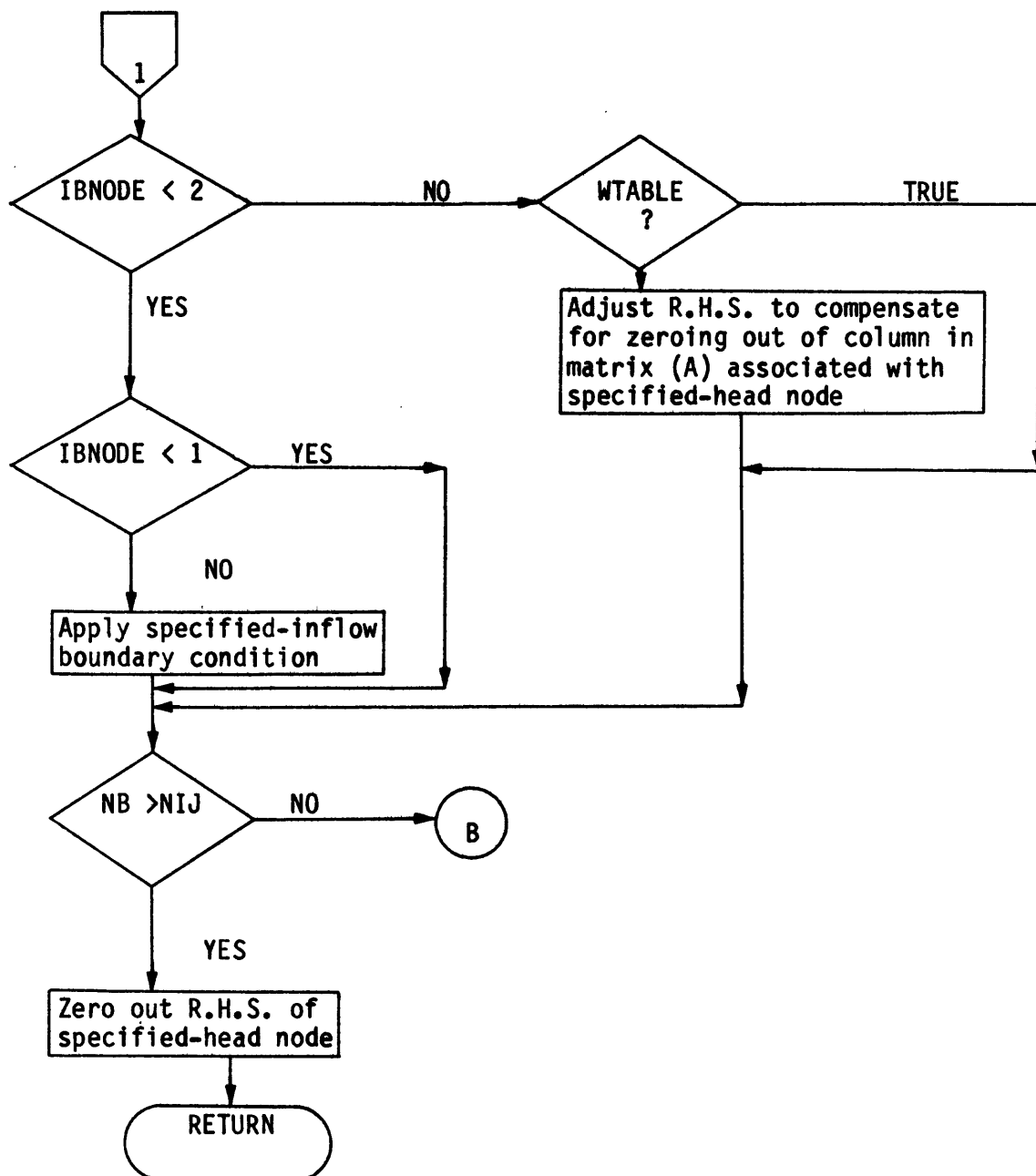


SUBROUTINE BNDCON

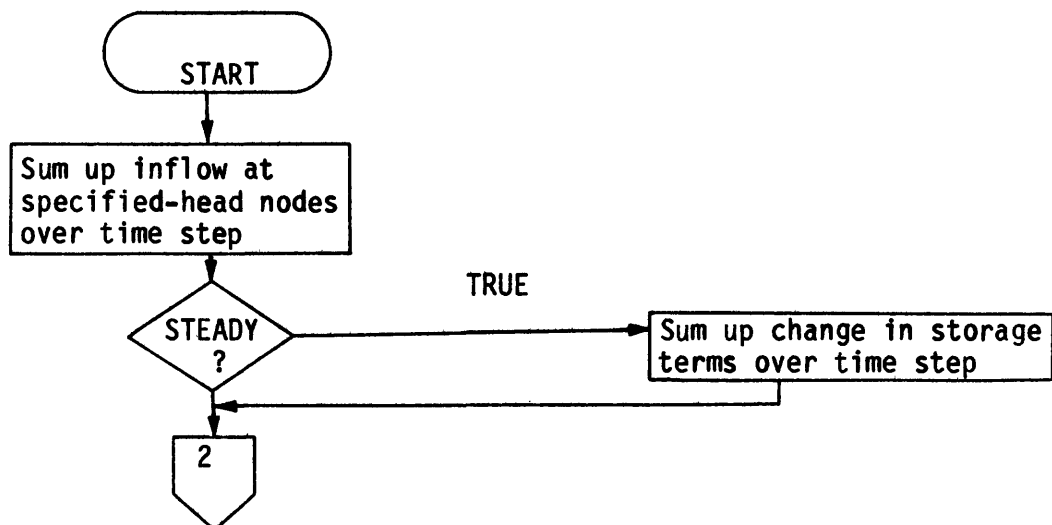


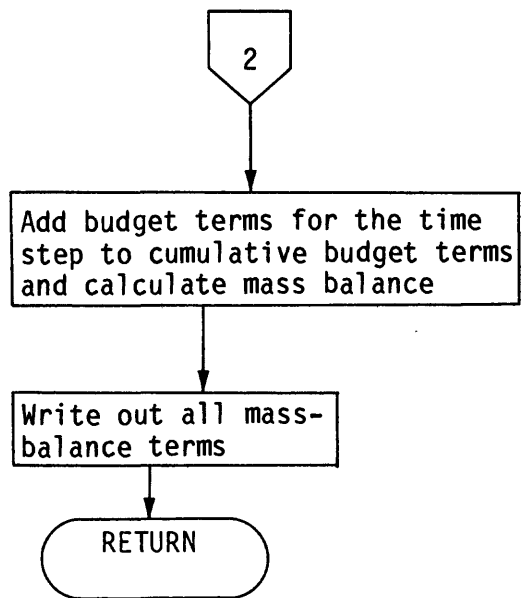
SUBROUTINE QBOUND



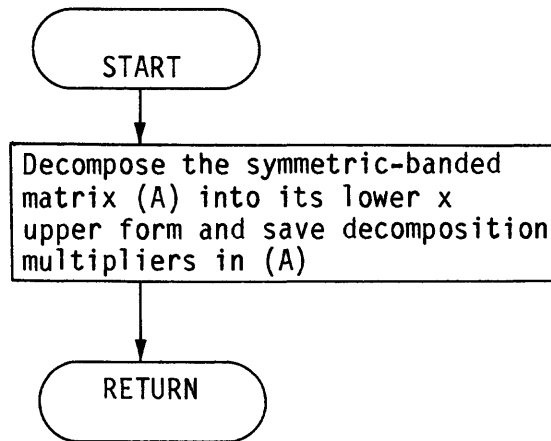


SUBROUTINE MASBAL

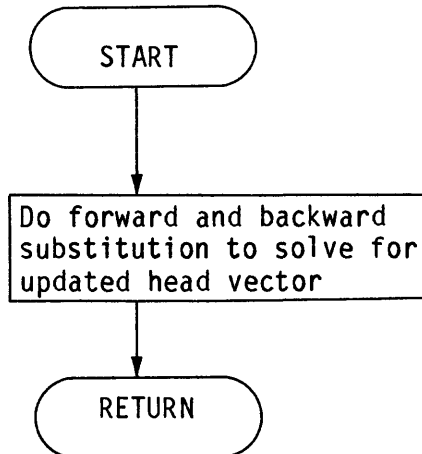




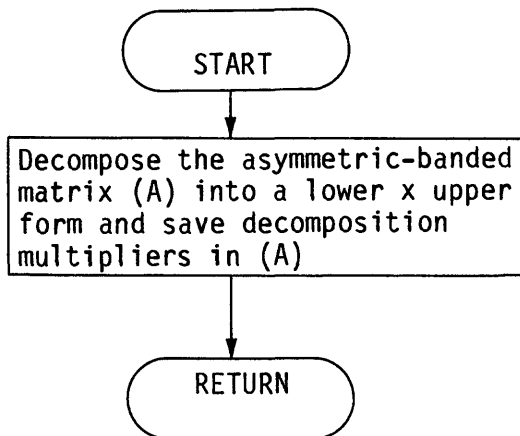
SUBROUTINE SDCOMP



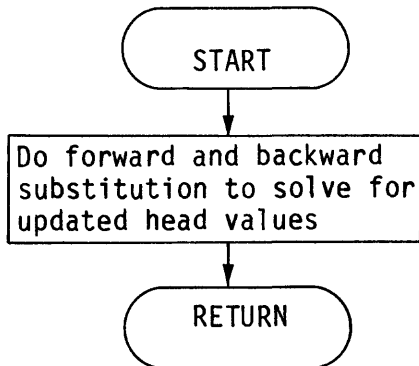
SUBROUTINE SSOLVE



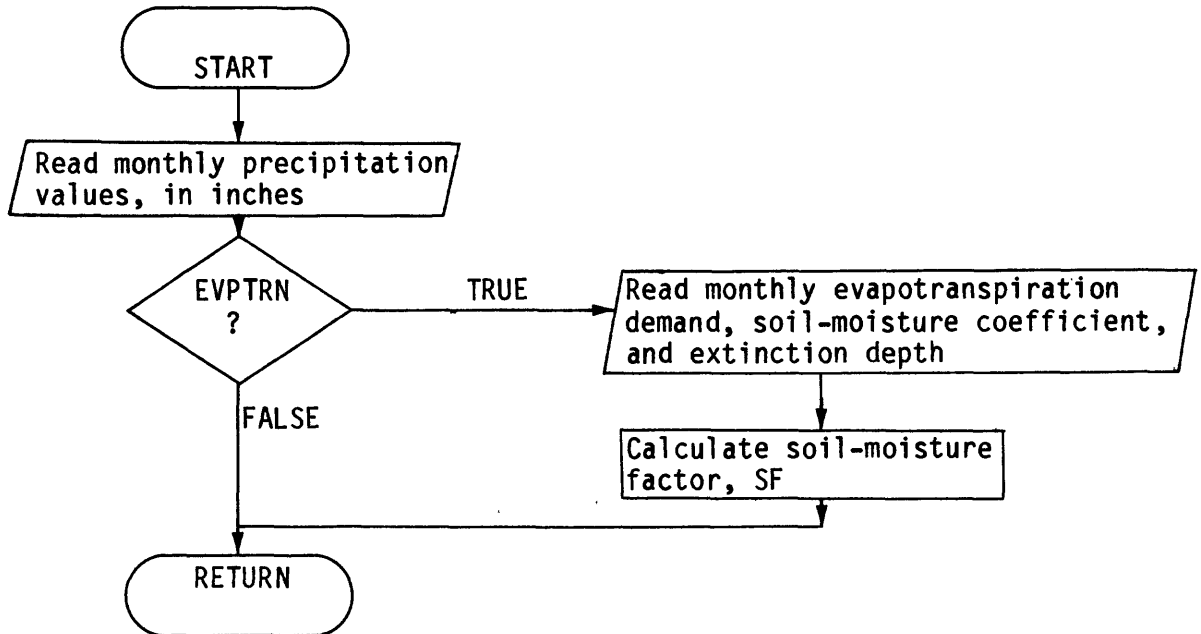
SUBROUTINE ADCOMP



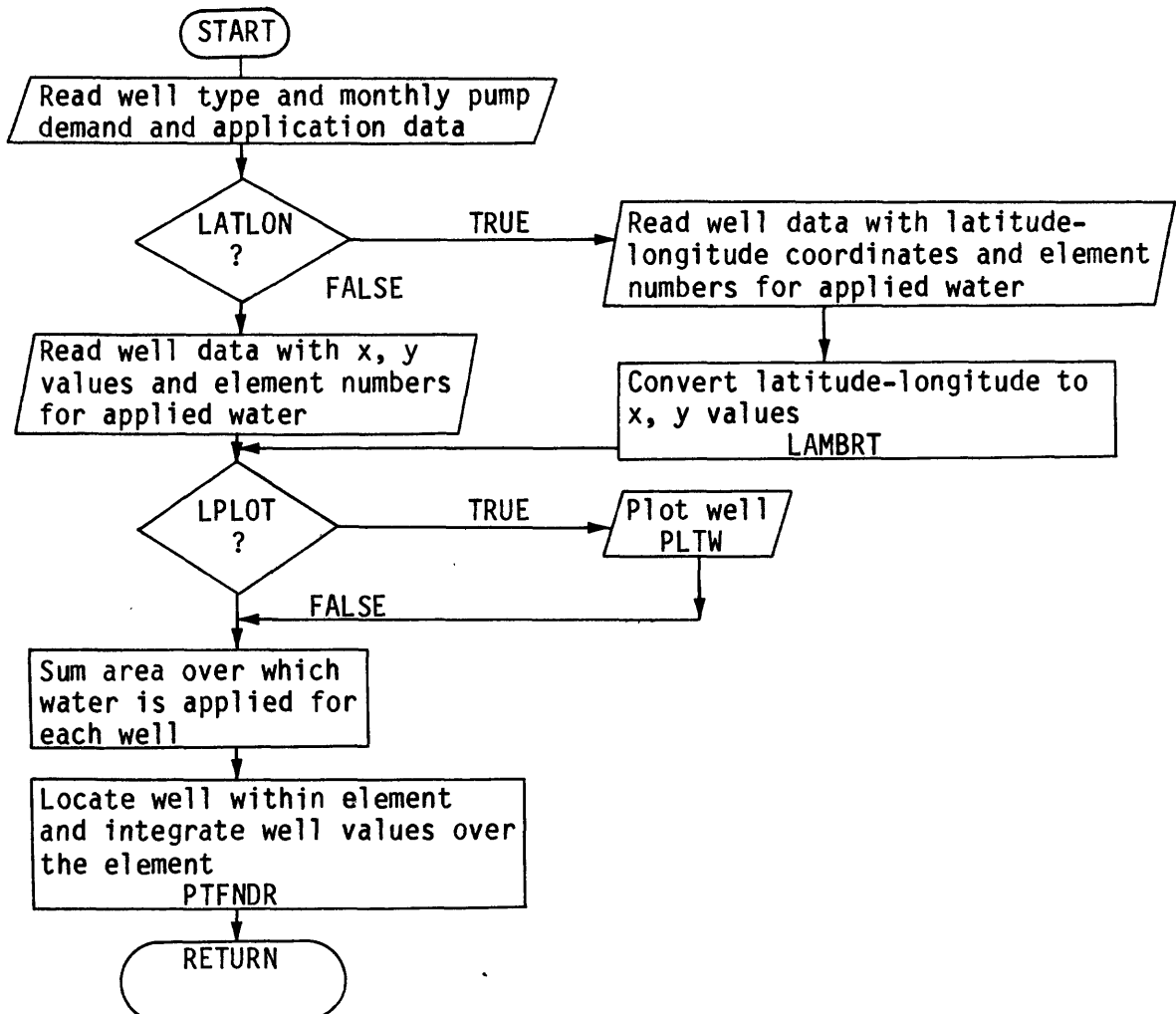
SUBROUTINE ASOLVE



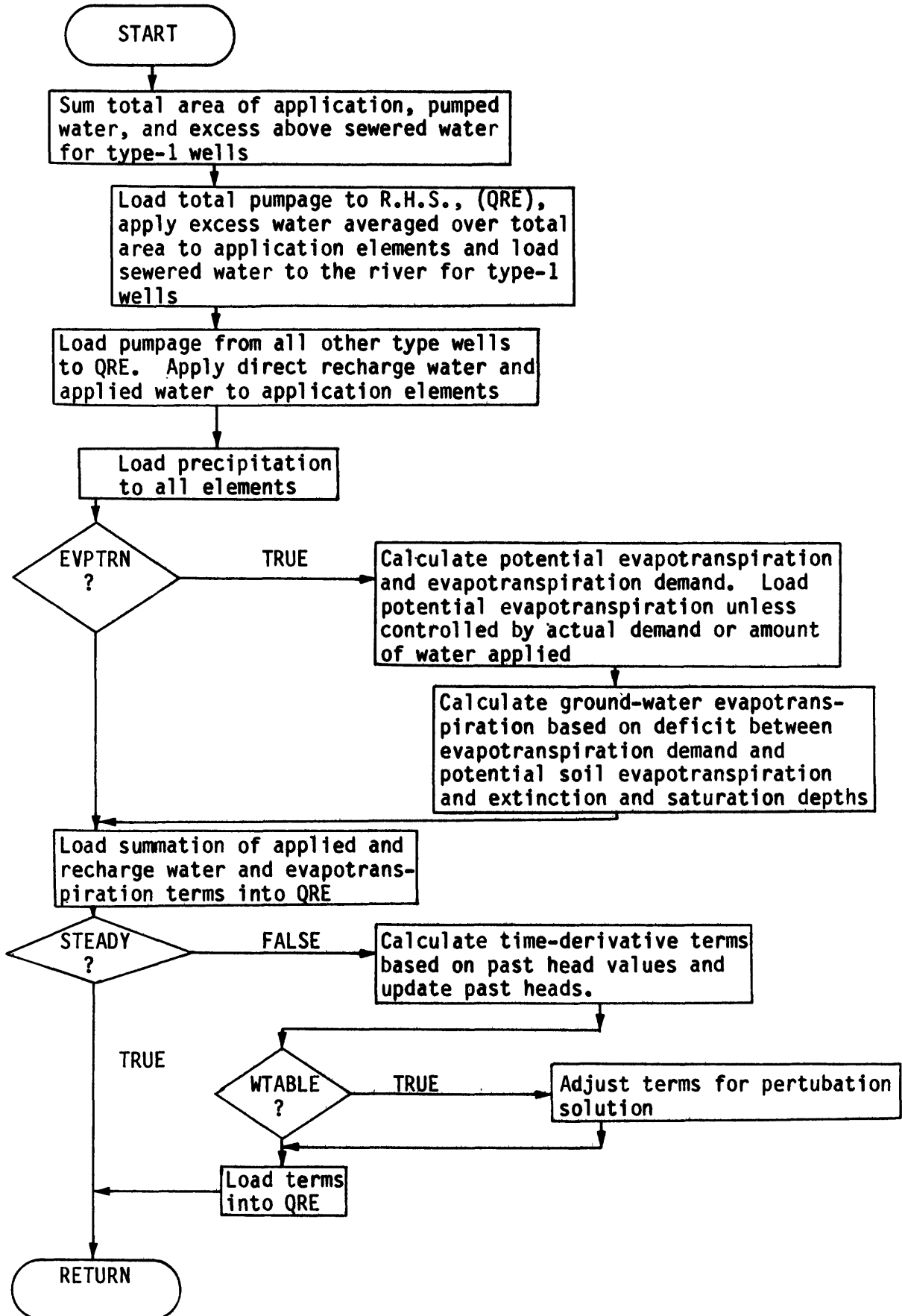
SUBROUTINE QZEE



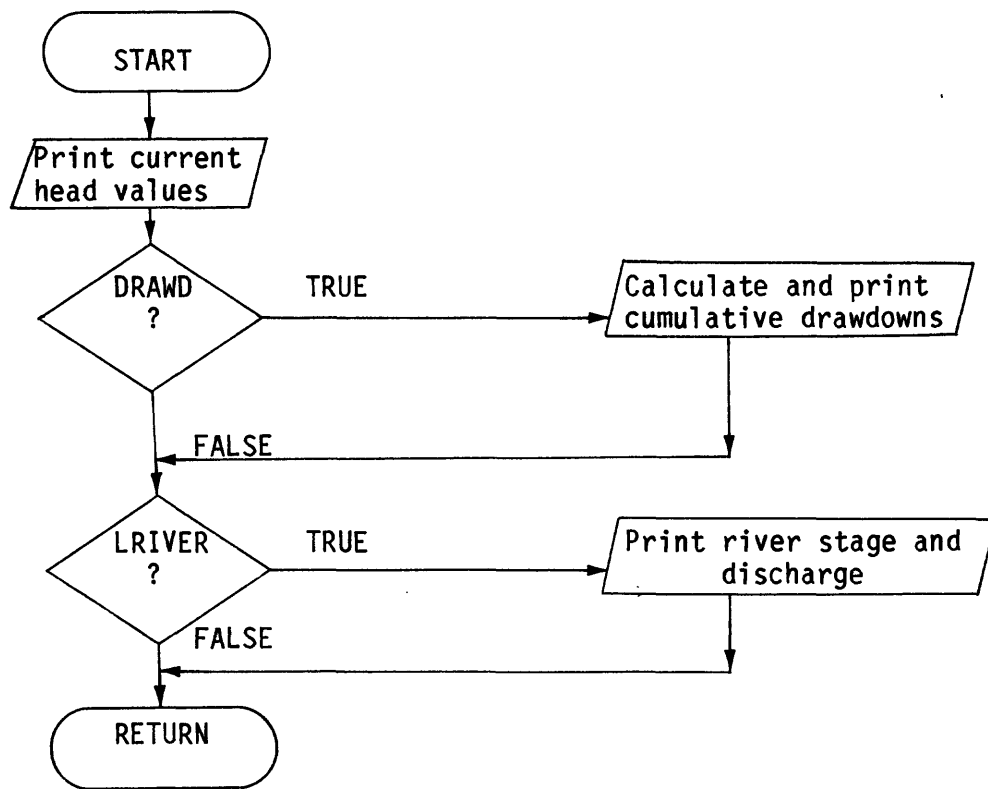
SUBROUTINE PUMPNU



SUBROUTINE LOADQR



SUBROUTINE OUTPUT



Model Input Setup

Group I: Title, Model Options, Printout and Plotter

Control, and Problem Size

This group of cards is read in by the DATAIN subroutine. Integer variables, read in with I format, must be right-hand justified. Logical variables are read in with L format and must be left-hand justified. Real variables are read in with F format, but they can also be punched as E format (for example, 0.00004 can also be punched as 4.0E-05).

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-72	9A8	TITLE	First title card used as a heading to be printed on the output
2	1-72	9A8	TITLE	Second title card
3	1-4	L4	STEADY	Logical variable to specify program options .TRUE. - steady-state simulation .FALSE. - transient run
	7-10	L4	WTABLE	.TRUE. - transmissivities vary with head (water-table simulation) .FALSE. - transmissivities are constant
	13-16	L4	QUAD	.TRUE. - elements are shaped as quadrilaterals .FALSE. - Elements are triangular
	19-22	L4	LATLON	.TRUE. - coordinates for all points are given by their latitude and longitude .FALSE. - points are defined by their x, y value
	25-28	L4	LRIVER	.TRUE. - calculations will be done to account for stream-aquifer interaction .FALSE. - no stream calculations
	31-34	L4	DRAWD	.TRUE. - drawdown values from the starting heads will be calculated

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
	37-40	L4	LPLOT	.FALSE. - only heads will be calculated .TRUE. - Plotting routines will be called
				.FALSE. - No plotting will be done
	43-46	L4	EVPTRN	.TRUE. - Evapotranspiration will be accounted for .FALSE. - no evapotranspiration

Note: If option space is left blank, it will be interpreted as .FALSE.

4	1-4	14	IPRINT	Printout option control IPRINT \geq 6, All nodal x, y coordinates are printed out IPRINT \geq 5, Nodal linkages and element areas are printed IPRINT \geq 4, River and well coordinates are printed IPRINT \geq 3, All input data are echoed IPRINT \geq 2, Critical input is echoed IPRINT \geq 1, Minimum printout
	5-8	14	NPRNT	Number of time steps between printout of calculated heads and river stage
5	1-4	14	KYR	Starting year of simulation
	5-8	14	KMN	Starting month
	9-12	14	KDY	Starting day of month
	13-16	14	IDAY	Time-step, in days
	17-20	14	NSTEP	Total number of time steps to be taken
6	1-4	14	MAXIT	Maximum number of iterations allowed in river-stage head calculation
	5-14	F10.0	BETA	Convergence acceleration parameter (generally equal to 1.0 or greater)
	15-24	F10.0	EPS	Convergence criterion between iterations

NOTE: CARD 6 read only if LRIVER = .TRUE.

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
7	1-6	F6.0	XSCL	Scaling factor for x coordinates to convert them to model-length units
	7-12	F6.0	YSCL	Scaling factor for y coordinates
NOTE: CARD 7 is read only if LATLON = .FALSE.				
8	1-4	14	IPLLOT	Control parameter for plot of the finite-element grid
				IPLLOT \geq 5, Nodes are numbered
				IPLLOT \geq 4, Elements are numbered
				IPLLOT \geq 3, River nodes are numbered
				IPLLOT \geq 2, Wells are numbered
				IPLLOT \geq 1, No numbering
	5-14	F10.0	XSCLP	Scaling factor to convert the x-coordinate values back to plotter inches
	15-24	F10.0	YSCLP	Scaling factor for y coordinates
NOTE: CARD 8 is read only if LPLLOT = .TRUE.				
9	1-4	14	IS	Number of nodes in the row (long dimension)
	58	14	JT	Number of nodes in the column (short dimension)

Group II: Array Data for the Aquifer

Each of the following data sets, except the first (IBNODE) and last two (PBH and PBF), are read in by the subroutine ARRAY and consist of a parameter card, and if the values are to be specified at each node, a set of data cards. All arrays (except the first) are real. Each parameter card has the following format:

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-8	A6	TITLE	Data-set name
	11-13	12	NUNIT	Unit number from which additional data cards will be read. If NUNIT = 0, all nodes are set to a uniform value, and no additional data are read

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
	14-29	2A6	FMT	A character string describing the format to be used in the subsequent data set [for example, (8F10.5)]
	30-39	F10.0	SCALE	Scaling factor for values read in the data set. If NUNIT = 0, the entire array is set to SCALE
	40-43	14	IPRINT	Local printout control. If IPRNT > 0, scaled values will be printed out

The arrays needed are determined by the logical control options specified. Those that are always read in are underlined. Those that are optional will be noted. Data-set one (IBNODE) is integer and read in by DATAIN. No parameter card is necessary. All data sets have IS number of cards with JT values per card.

<u>Data card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-80	2014	IBNODE(I,J)	Integer describing boundary-condition type at each node IBNODE (I,J) = 0, Interior node with no specified influx IBNODE (I,J) = 1, Specified-flux type boundary IBNODE (I,J) = 2, Specified-head type boundary IS cards, JT values per card
2	1-80	FMT	<u>H(I,J)</u>	Starting nodal head values [L]
3	1-80	FMT	HP(I,J)	Head values at the last time step in a previous run [L] NOTE: Read in when STEADY = .FALSE.
4	1-80	FMT	S(I,J)	Nodal values for the storage coefficient [dimensionless] NOTE: Read in when STEADY = .FALSE. Data sets 3 and 4 are needed only for a transient simulation

<u>Data card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
5	1-80	FMT	<u>T(I,J)</u>	<p>If WTABLE = .TRUE., the array contains the nodal values of permeabilities [L/T]</p> <p>If WTABLE = .FALSE., the array contains the nodal values of transmissivities [L]</p>
6	1-80	FMT	ALSD(I,J)	<p>Elevation of the land surface above the datum used in calculating evapotranspiration [L]</p> <p>Note: Read in when EVPTRN = .TRUE.</p>
7	1-80	FMT	HO(I,J)	<p>Steady-state solution values for head, used in the perturbation-solution technique for variable transmissivities [L]</p> <p>Note: Read in when WTABLE = .TRUE.</p>
8	1-80	FMT	BOTT(I,J)	<p>Elevation of the aquifer bottom above the datum [L]</p> <p>Note: Read in when WTABLE = .TRUE.</p> <p>Data sets 7 and 8 are needed when doing a water-table aquifer simulation</p>
9	1-80	FMT	<u>BQH(I,J)</u>	<p>Array of values for specified head and fluxes at boundary nodes [L] or [L³/T]</p> <p>Note: Boundary-condition type is determined by the IBNODE array</p>
10	1-72	12F6.0	<u>PBH(M)</u>	<p>Monthly multipliers for specified-head values. One card only. No parameter card needed.</p>
11	1-72	12F6.0	<u>PBF(M)</u>	<p>Monthly multipliers for specified flux values. One card only. No parameter card needed.</p>

Group III: Finite-Element Grid Specifications

This data set contains the information necessary to generate the finite-element grid. Because the grid is regular, only the coordinates of the top and bottom nodes on a line need be specified. JT number of nodes will then be spaced equally along that line. The coordinates can be located by their latitude and longitude or by user determined x, y value. Both formats will be detailed below. The grid specification is the same whether the elements are quadrilaterals or triangular.

NOTE: if LATLON = .TRUE., data are read with the following format.
There are IS number of cards in the data set.

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1-IS	1-12	314	LONB	Values for degrees, minutes, and seconds of longitude for a node in the bottom row of the grid (left to right). Each value must be right-hand justified in its 14 field
	13-24	314	LATB	Degrees, minutes, and seconds of latitude for nodes in the bottom row
	31-42	314	LONT	Degrees, minutes, and seconds of longitude for corresponding nodes in the top row
	43-54	314	LATT	Degrees, minutes, and seconds of latitude for nodes in the top row

Note: if LATLON = .FALSE., data are read with the following format; values will be scaled by XSCL and YSCL.

1-IS	1-10	F10.0	X(NODBOT)	x coordinate of the bottom node [L]
	11-20	F10.0	Y(NODBOT)	y coordinate of the bottom node [L]
	21-30	F10.0	X(NODTOP)	x coordinate of the top node [L]
	31-40	F10.0	Y(NODTOP)	y coordinate of the top node [L]

Group IV: River Geometry, Leakance Values, Stage-Discharge Relationships, and Daily Inflow Values

This group contains data necessary to account for stream-aquifer interaction. This group is read only if LRIVER = .TRUE. and is read by subroutine QRIVER.

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-4	14	NREACH	Number of reaches into which the river is divided

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
	5-8	14	NRNODE	Number of river nodes
	9-12	14	NRZ	Number of zones defined by river-leakance values
	13-16	14	NTRIB	Number of river tributaries
2	1-72	7F10.0	QCOR(1)	Leakance (K'/b') values for each zone [1/T]. There are NRZ values, 12 values per card

There are three data sets to be read in the river-data group. The first set contains data for river-flow routing, including pool-elevation data and the stage-discharge relationship for each reach. There are three cards for each reach and NREACH groups of cards in the set. Data-set two contains river-node parameters, and there are NRNODE cards. Data-set three contains the daily inflow values at the upstream end of the river.

DATA-SET ONE: Pool-elevation and stage-discharge data

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-4	14	NSTAGE(N)	Number of points used to approximate the stage-discharge relationship for this reach
	5-14	F10.0	HPOOL(N)	Pool elevation for this reach [L]
2	1-72	12F6.0	STAGE(L,N)	Values for the stage at selected points along the stage-discharge curve [L]. There are NSTAGE(N) values, 12 values per card
3	1-72	12F6.0	DSCHRG(L,N)	Corresponding values for discharge at selected points along the S-D curve [L^3/T]

DATA-SET TWO: River-node parameters

NOTE: If LATLON = .TRUE., data are read with the following format.

NRNODE cards

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-12	314	LON	Degrees, minutes, and seconds of longitude for each river node
	13-24	314	LAT	Degrees, minutes, and seconds of latitude for each river node

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
	25-28	14	NRN(N)	Reach number of the node
	29-32	14	IRZ	Leakance-zone number of the node
	33-36	14	NER(N)	Number of the element in which the river node is located
	37-40	14	NRTYP(N)	River-node type NRTYP(N) = 0, stage-discharge control node NRTYP(N) = 1, pool-backwater control node
	41-50	F10.0	LENGTH	Channel length associated with the river node [L]
	51-60	F10.0	WIDTH	Channel width at the river node [L]
	61-70	F10.0	ELVR(N)	Elevation of the riverbed above the datum at the river node [L]
	71-80	F10.0	XFACTR(N)	NRTYP(N) = 0, not used NRTYP(N) = 1, interpolation factor to determine stage from pool and previous river-node stage elevation

NOTE: If LATLON = .FALSE., data is read in with the following format;
x and y values are scaled by XSCL and YSCL.

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-12	F12.0	XR	x coordinate of the river node [L]
	13-24	F12.0	YR	y coordinate of the river node [L]
	25-28	14	NRN(N)	Reach number of river node
	28-32	14	IRZ	Leakance-zone number
	33-36	14	NER(N)	Element number of river node
	37-40	14	NRTYP(N)	River-node type
	41-50	F10.0	LENGTH	Channel length [L]
	51-60	F10.0	WIDTH	Channel width [L]

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
	61-70	F10.0	ELVR	River elevation
	71-80	F10.0	XFACTOR(N)	Interpolation factor
2	1-80	20I4	NODEIN (NTRIB)	River-node numbers of tributary-input point

DATA-SET THREE: Daily inflow values

1-31	1-72	12F10.0	FLOWIN(N)	Daily values for inflow at the upstream end of the river [L ³ /T]. 365 or 366 values, 12 values per card
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Group V: Precipitation and Evapotranspiration Data

This set of cards contains data used in calculating recharge due to precipitation and is read by subroutine QZEE. If EVPTRN = .TRUE., data used in calculating losses due to evapotranspiration are also read. Precipitation, evapotranspiration demand, and soil-moisture capacity are read with units of inches. NOTE: The first value on the card is assumed to be for the month of January.

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-72	12F6.0	PRECIP(M)	Monthly averaged value for precipitation [L]
2	1-72	12F6.0	U(M)	Monthly averaged value for evapotranspiration demand [L]
3	1-6	F6.0	SC	Soil-moisture capacity [L]
	7-12	F6.0	DEPTH	Extinction depth for soil-moisture loss to evapotranspiration [L]

NOTE: Read only if EVPTRN = .TRUE.

Group VI: Well Parameters, Pumpage Demands, and Water Distribution

This group contains data needed to calculate the effects due to groundwater withdrawal and reapplication. Wells can be typed as to percentage of water that is applied in irrigation and seasonal pumpage variation. Sewered wells (type 1) are defined as those whose discharge is routed directly into the river (for example, municipal wells).

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-4	14	NWELLS	Total number of wells
	5-8	14	NSEWER	Number of type-1 wells whose outflow contributes directly to the river
	9-12	14	NWTYP	Total number of well types

There are two additional data sets to be read. The first describes the monthly variation in the pump demand by type and the amount of water applied. The second set contains the well location, type, and withdrawal rate and which elements, if any, the water is applied over. Both data sets are read only if NWELLS > 0.

DATA-SET ONE: Monthly variation in pump demand and percentage applied

1	1-60	12F5.0	PMPDMD(N,M)	Monthly multipliers of pump demand by pumpage types. There are NWTYP cards in this data set.
	61-65	F5.0	PMPAPP(N)	Percentage of well water that is applied on land by pumpage types.

DATA-SET TWO: Well parameters

NOTE: If LATLON = .TRUE., the data set is read with the following format.

There are NWELLS sets of cards.

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
1	1-12	314	LON	Degrees, minutes, and seconds of longitude for each well
	13-24	314	LAT	Degrees, minutes, and seconds of latitude for each well
	25-28	14	NEW(N)	The number of the element in which the well is situated
	29-32	14	NTW(N)	Well-type number
	33-36	14	NAPP(N)	The number of elements over which the applied water is distributed
	41-50	F10.0	QW(N)	Pumpage rate from the well [L ³ /T]. <u>For recharge wells, QW(N) must be negative</u>

<u>Card</u>	<u>Columns</u>	<u>Format</u>	<u>Variable</u>	<u>Definition</u>
2	1-80	20I4	NEAPP(L,N)	Element numbers on which water is applied. Twenty values per card NOTE: This card is read only if NAPP(N) for the well is >0
NOTE: If LATLON = .FALSE., the data set is read with the following format. x- and y-coordinate values are scaled by XSCL and YSCL				
1	1-12	F12.0	XW	x coordinate of the well [L]
	13-24	F12.0	YW	y coordinate of the well [L]
	25-28	14	NEW(D)	Number of the element containing the well
	29-32	14	NTW(N)	Well-type number
	33-36	14	NAPP(N)	Number of elements over which water is applied
	41-50	F10.0	QW(N)	Pumpage rate [L ³ /T]
2	1-80	20I4	NEAPP(L,N)	Element numbers on which water is applied. Twenty values per card NOTE: This card is read only if NAPP(N) for the well is >0.

Model Input for Example Problem

A 20 x 20 node grid (fig. 28) is used to illustrate the model-input setup. Grid spacing is 1,000 feet. Model options included in the example problem are confined-aquifer conditions, precipitation, evapotranspiration, river, one injection well, one pumping well, constant-head boundary, and constant-flux boundary.

Aquifer and model properties are as follows:

Initial potentiometric surface	200 feet
Storage coefficient	0.01
Transmissivity	10,000 ft ² /d
Well injection rate	8.64 x 10 ⁴ ft ³ /d
Well pumping rate	1.34 x 10 ⁵ ft ³ /d
River: 18 nodes	
Bed elevation	199 feet
Stage-discharge:	
stage	0.0 1.0 2.0 feet
discharge	0.0 1.0 100.0 ft ³ /s
Reach length	1,000 feet
Reach width	40 feet
Leakage	0.15 day ⁻¹
Daily inflows	50 ft ³ /s

Precipitation (January)	0.43 inch per month
Evapotranspiration demand (January)	0.49 inch per month
Soil-moisture storage capacity	6 inches
Extinction depth for evapotranspiration	10.0 feet
Constant-head boundary	200 feet
Constant-flux boundary	0.01 ft ³ /s

A listing of model input (p. 160-162) and a partial listing of model output (p. 163-168) for the example problem shown in figure 28 are included at the end of this report. Line numbers are shown to facilitate reading the example but should not be included in actual model input. Also note that the model is presently setup to print 60 river nodes in pairs; however, only the nodes pertinent to the problem were retained in the example output.

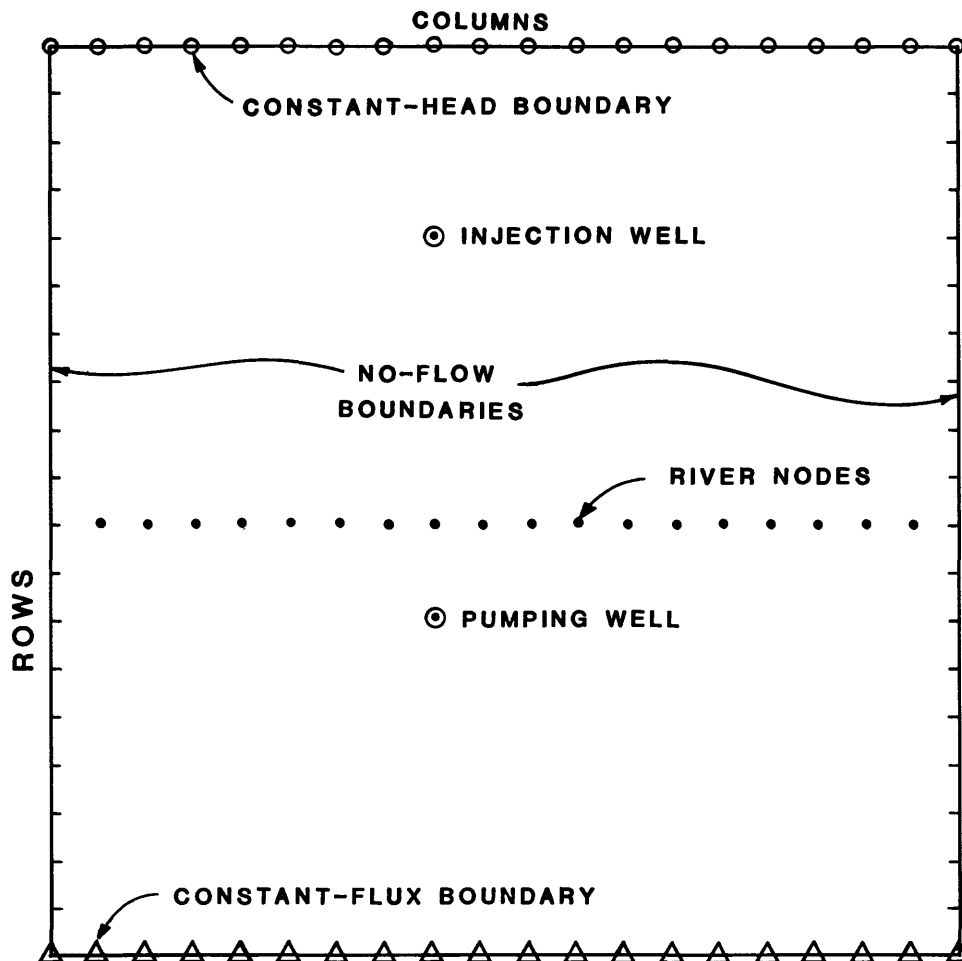


Figure 28.--Model characteristics for input example problem.

Description of Program Common Blocks

- AQUIFER - Contains the element area array and all arrays associated with aquifer properties.
- BNDRY - Contains all arrays used in applying model-boundary conditions.
- BUDGET - Contains arrays used in computing the water balance for the model.
- HEAD2 - Contains arrays to store original and iterative head values.
- IOUNIT - Contains the variables defining the input and output units and the variables used as output control.
- LOGIC - Contains the logical variables controlling program operation.
- MATRIX - Contains the matrix used as the left-hand side of the flow equation and the matrix containing the time-derivative terms.
- PLT - Contains variables used in the grid-plotting routines.
- RIVERB - Contains all data and arrays associated with the stream being modeled.
- SIZE - Contains all variables that detail the dimensions of the problem and that describe the element geometry.
- TIME - Contains data used in keeping track of calendar time for transient simulations.
- TIME2 - Contains control variables for the time and iteration loops.
- WELLS - Contains all data relating to well pumpage and water application.
- WETHER - Contains climate and evapotranspiration data.
- WTBLE - Contains additional arrays used when transmissivities are head dependent (water-table simulation).

Array Dimensions by Common Blocks

Minimum dimensions for each array is given in terms of the integer variable that controls its size. Where a number is given, it indicates that the array dimension is fixed and never needs to be changed unless otherwise noted. More information about dimensioning of variables is given in the list of variable definitions.

Common Block

Variables

AQUIFER:	ALSD (NIJ)	QZ (NIJ)
	AREA (NEL)	S (NIJ)
	H (NIJ)	T (NIJ)
	HP (NIJ)	X (NIJ)
	QRE (NIJ)	Y (NIJ)
BNDRY:	BB (9,50)	PBF (12)
	BQH (NIJ)	PBH (12)
	IBNODE (NIJ)	QB (NIJ)
BUDGET:	A (NIJ, NBW)	NGL (4, NEL)
	BUD (NBUD)	NRN (NRNODE)
	BUDTOT (NBUD)	NRTYP (NRNODE)
	BUDTTL (2,NBUD)	NSTAGE (NREACH)
	DISCHRG (20,NREACH)	OLDH (NIJ)
	ELVR (NRNODE)	QCOMR (NRNODE)
	FLOWIN (366)	QCOR (NRZ)
	FMON (12)	RIVERH (NRNODE)
	HI (NIJ)	RIVERQ (NRNODE)
	HPOOL (NREACH)	RW (4,NRNODE)
	IDMON (12)	STAGE (20,NREACH)
	MON (12)	UV (NIJ)
	NEND (NREACH)	XFACTR (NRNODE)
	NER (NRNODE)	
WELLS:	AREAPP (NWELLS)	PMPAPP (NWTYP)
	ETFACT (8)	PMPDMD (NWTYP,12)
	NAPP (WELLS)	QAPP (NEL)
	NEAPP (30,NWELLS)	QW (NWELLS)
	NEW (NWELLS)	REAPP (NEL)
	NTW (NWELLS)	U (12)
	PRECIP (12)	WW (4,NWELLS)
WTBLE:	BOTT (NIJ)	
	HO (NIJ)	
	QBAR (NIJ)	

Definitions of Program Variables

A(NIJ,NBW)	- The banded matrix containing all terms on the left-hand side of the flow equation. When transmissivities are constant, the A matrix is symmetric, and only the upper half of the matrix is stored. When transmissivities are head dependent, the matrix is asymmetric, and the full band is stored.
ALSD(NIJ)	- Array of land-surface elevations.
AREA(NEL)	- Array of element areas.
AREAPP(NWELLS)	- Total area of water application for each well.
BB(9,50)	- Storage for full matrix equations at specified-head boundary nodes. The "9" is fixed, but the "50" should be increased if there are more than 50 specified-head nodes.
BETA	- Acceleration or dampening factor used in controlling the iterative solution for river leakage.
BOTT(NIJ)	- Elevations of the aquifer bottom measured at the nodes.
BQH(NIJ)	- Boundary values of specified heads and fluxes.
BUD(NBUD)	- Time-step flow rates of each budget item.
BUDTOT(NBUD)	- Accumulated volume of flow for each budget item.
BUDTTL(2,NBUD)	- Names of budget (water-balance) items. Size of array may be machine word-size dependent.
DELT	- Size of the time step in seconds.
DEPTH	- Evapotranspiration extinction depth (feet).
DRAWD	- Controls whether cumulative drawdowns from starting head values will be calculated. .TRUE. for calculations to be done.
DSCHRG(20,NREACH)	- Corresponding river-reach gage discharge of stage-discharge relationship.
ELVR(NRNODE)	- Nodal river-bottom elevation array.
EPS	- Convergence criterion for iterative procedure.

ETFACT(8)	-	Coefficients used in calculating evapotranspiration from demand and moisture capacity.
EVPTRN	-	Controls whether evapotranspiration will be considered. .TRUE. for evapotranspiration calculation.
FLOWIN(366)	-	Daily values for inflow at the first upstream river node.
FMON(12)	-	Percentages of days out of the time step that are spent each month.
FMT	-	The format with which the subsequent data will be read (NUNIT >0).
GRADX	-	Mean gradient in x-direction of 'H0' over an element.
GRADYY	-	Mean gradient in y-direction of 'H0' over an element.
H(NIJ)	-	Current value of calculated heads at each node.
HI(NIJ)	-	Array for storing head values at the current iteration level.
HO(NIJ)	-	The array of head values defining an approximate average head surface <u>that satisfies the ground-water-flow equation.</u>
HP(NIJ)	-	Calculated value of heads at the previous time step.
HPOOL(NREACH)	-	River-reach pool elevations.
IBNODE(NIJ)	-	Integer array defining boundary type at each node.
IDAY	-	Number of days per time step.
IDIAG	-	The location, within the A matrix, of the diagonal. (IDIAG = 1, for constant transmissivity; for variable transmissivity, IDIAG = JT + 2).
IDMON(12)	-	Array containing the number of days in each month.
IDYR	-	Total number of days in the year. Equal to 365 or 366.
IHALFB	-	The half-band width of matrix A, not including diagonal.

IHBP	-	The half-band width of matrix A, including diagonal.
IN	-	Unit number of input device.
IO	-	Unit number of output device. Both IN and IO may be machine dependent and are defined in BLOCK DATA.
IPLOT	-	Integer controlling the detail of the grid plot.
IPRINT	-	Printout control variable.
IS	-	An integer defining the number of nodes along the length of the finite-element grid.
JT	-	The number of nodes along the width of the grid.
KDY	-	Calendar day for transient simulation.
KDYS	-	Number of days from the beginning of the calendar year to the current calendar date in the simulation.
KMN	-	Calendar month for transient simulation.
KYR	-	Calendar year for transient simulation.
LATLON	-	Determines whether all points are specified by latitude and longitude or by x, y values. .TRUE. for latitude and longitude.
LPLOT	-	Determines whether plot routines will be called. .TRUE. for plotting.
LRIVER	-	Controls whether stream-aquifer interaction will be evaluated. .TRUE. if there is a contributing stream.
MAXIT	-	Maximum number of iterations allowed in river-leakage procedure.
MON(12)	-	Array of three-letter abbreviations of the month names.
NAPP(WELLS)	-	Array containing the total number of elements over which water is applied for each well.
NBUD	-	Number of budget items. Defined by the number of river reaches (NREACH) plus 12.

NBW	-	The matrix half-band width plus the diagonal (JT + 2) for the constant transmissivity case. The full-band width $[2(JT + 2) - 1]$ for the variable transmissivity problem.
NE	-	The element that the point is located within.
NEAPP(30, NWELLS)	-	Array containing the element numbers over which water is applied. Up to 30 elements are allowed per well.
NEL	-	Total number of elements the grid is composed of. $[(IS-1) \times (JT-1)]$ for quadrilateral elements. Twice that amount for triangles].
NEND(NREACH)	-	The river-node number defining the downstream end of each reach.
NER(NRNODE)	-	An array containing the element in which each river node is located.
NETYP	-	An integer defining the number of nodes in the elements. (NETYP = 3 for triangular elements; for quadrilaterals, NETYP = 4).
NEW(NWELLS)	-	Array containing the element number in which each well is situated.
NGL(4,NEL)	-	The array defining the nodal linkages of each element. Nodes are linked counterclockwise around the element.
NIJ	-	The total number of nodes (IS x JT).
NO	-	The well or river-node number that the point represents.
NPRNT	-	Number of time steps between output of heads and river stages.
NREACH	-	An integer declaring the number of reaches the river is sectioned into.
NRN(NRNODE)	-	An array containing the reach number for all river nodes.
NRNODE	-	The number of river nodes.
NRTYP(NRNODE)	-	This array contains integer values declaring the node type for each river node. If NRTYP(N) = 0, the node is stage-discharge controlled. If NRTYP(N) = 1, it is pool-backwater controlled.

NRZ	-	The number of zones the river is divided into based on the value for leakance.
NSEWER	-	The number of type-1 wells.
NSTAGE(NREACH)	-	An array of integers defining the number of points used in describing the linearized stage-discharge relationship for each reach.
NSTEP	-	Total number of time steps to be taken in transient simulation.
NTW(NWELLS)	-	Integer-defining well type for each well. Well types are defined by user as to amount of applied water and seasonal variability. Type-1 wells, however, are defined as those whose excess water contributes directly to streamflow.
NUNIT	-	The unit number from which the subsequent data will be read. If NUNIT = 0, all nodes will be set to a uniform value.
NWELLS	-	The total number of wells.
NWTYP	-	The total number of well types.
OLDH(NIJ)	-	Array for storing starting head values.
PBF(12)	-	Multipliers for specified-flux values.
PBH(12)	-	Multipliers for specified-head values to account for monthly variations.
PIVOT	-	The diagonal term of each row.
PMPAPP(NWTYP)	-	Percentage of pumped water that is applied by well type.
PMPDMD(NWTYP,12)	-	Monthly adjustment values for pumpage rates by the type of well.
PRECIP(12)	-	Monthly averaged precipitation (inches).
QAPP(NEL)	-	Array of applied-water values that are subject to evapotranspiration losses averaged over the elements.
QB(NIJ)	-	Total fluxes to boundary nodes.
QBAR(NIJ)	-	The array of flux values that sustain the H0 head surface.

QCOMR(NRNODE)	-	River-node leakage coefficient, equal to leakage factor of the zone the node is in times the river width and length associated with the node.
QCOR(NRZ)	-	The array of leakance values defined by leakance zones.
QRE(NIJ)	-	Nodal-recharge array without explicit river-leakage terms.
QUAD	-	Determines the type of finite element used. .TRUE. if quadrilateral, .FALSE. for triangles.
QW(NWELLS)	-	Annual average pumping rate at each well, in cubic feet per second.
QZ(NIJ)	-	Total nodal recharge array.
RADI	-	Conversion factor for degrees to radians.
RAIN	-	Monthly precipitation value, in feet per second.
REAPP(NEL)	-	Array of applied water that is direct recharge not subject to evapotranspiration, averaged over the elements.
RIVER(NRNODE)	-	Array for storing river-node stage elevations.
RIVERQ(NRNODE)	-	Array for storing nodal river-flow values.
RLATA	-	The latitude of origin of the Lambert projection.
RLONGA	-	The longitude of the origin of the Lambert projection.
RW(4,NRNODE)	-	Nodal values of leakage distribution from river nodes to element nodes.
S(NIJ)	-	Nodal storage array.
SC	-	Soil-column moisture capacity (inches).
SCALE	-	The scaling factor by which all nodes will be multiplied.
SEWAGE	-	Accumulated volume of pumped water to be routed downstream.
SF	-	Actual evapotranspiration.
STAGE(20,NREACH)	-	River-reach gage stage of stage-discharge relationship. Twenty points are allowed.

STEADY	-	Logical variable controlling whether run is steady state or transient. .TRUE. if steady state.
STOR	-	Mean storage term over an element.
T(NIJ)	-	Nodal transmissivity or permeability array.
TITLE	-	The name of the array of nodal values being read in.
TQH	-	Specified head or flux adjusted for monthly variability.
TRAN	-	Mean transmissivity over an element. (For a water-table aquifer transmissivity is average permeability times the thickness based on H).
U(12)	-	Monthly averaged evapotranspiration demand (inches).
UV(NIJ)	-	Storage for the time-derivative terms of the flow equation.
UET	-	Monthly values for evapotranspiration demand.
WTABLE	-	Controls whether transmissivities are constant or variable. .TRUE. if they are variable.
WW(4,NWELLS)	-	Distribution of pumpage from well points to element nodal points.
X(NIJ)	-	Array of nodal x coordinates.
XFACTOR(NRNODE)	-	Nodal value of nondimensional length from node to a downstream control or pool elevation.
XO	-	The x coordinate of the origin of the finite-element grid.
XP	-	The global coordinate of the point.
XR	-	The x coordinate of the origin of the Lambert projection.
XSCLP	-	Conversion factor for x coordinates back into plotter inches.
XW	-	Well x coordinate.
Y(NIJ)	-	Array of nodal y coordinates.

YO	-	The y coordinate of the origin of the finite-element grid.
YP	-	The global coordinate of the point.
YR	-	The y coordinate of the origin of the Lambert projection.
YSCLP	-	Conversion factor for y coordinates.
YW	-	Well y coordinate.

Tracy's Model Listing

C

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IMPLICIT DOUBLE PRECISION (A-H,O-Z)
REAL BUDTTL,XSCLP,YSCLP
LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON/AQUIFR/ H(572),HP(572),QRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
COMMON/WTBLE/ QBAR(572),BOTT(572),HO(572)
COMMON /BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
COMMON/BNDRY/BQH(572),QB(572),PBH(12),PBF(12),BB(9,50),IBNODE(572)MAI 20
COMMON/RIVERB/ HPOOL(2),STAGE(20,2),OSCHRG(20,2),QCOMR(118),MAI 40
1ELVR(118),XFACTR(118),RW(4,118),RIVERH(118),MAI 50
2RIVERQ(118),FLOWIN(366,7),QCOR(5),SEWAGE,NODEIN(7),NSTAGE(2),MAI 60
3NRN(118),NER(118),NRTYP(118),NREACH,NRNODE,NTRIB,QRR(64)MAI 70
COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYPMAI 80
COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KOYS,IDYR,MON(12)MAI 90
COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRNMAI 100
COMMON/WELLS/ QW(200),AREAPP(200),QAPP(1020),REAPP(1020),MAI 120*
1W(4,200),PMPDMD(5,12),PMPAPP(5),NEW(200),NTW(200),NAPP(200),MAI 130
2NEAPP(30,200),NWELLS,NSEWER,NWTPYMAI 140
COMMON /WETHER/ PRECIP(12),U(12),SC,DEPTH,SF,ETFAC(8)MAI 150
COMMON /PLT/ XSCLP,YSCLP,IPLOTMAI 160
COMMON /IOUNIT/ IN,IO,IPRINT,NPRNTMAI 170
COMMON/MATRIX/ A(572,50),UV(572)MAI 180
COMMON/HEAD2/ HI(572),OLDH(572)MAI 190
COMMON/TIME2/ DELT,BETA,EPS,NSTEP,MAXIT ,NNMAI 200
CHARACTER *20 JCNFIN,JCNFOUT,JCELPLT,EHEAD,JCPLOTDMAI 210
CHARACTER *20 JCRIVOUT,PORTOUT,STOROUTMAI 220
MAI 230
MAI 240
MAI 250

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C

```

1000 READ(1,1000)JCNFIN
      READ(1,1000)JCNFOUT
      READ(1,1000)JCELPLT
      FORMAT(A20)
      OPEN(UNIT=15,FILE= JCNFIN )
      OPEN(UNIT=16,FILE= JCNFOUT )
      OPEN(UNIT=18,FILE= JCELPLT )

```

C

C

```

C      READ IN INPUT DATA
C
C      CALL DATAIN
C
C      FORMULATE THE GLOBAL ELEMENT MATRICES
C
C      IF (QUAD) CALL LEMENQ (A,UV,NIJ,NBW)
C      IF (.NOT.QUAD) CALL LEMENT (A,UV,NIJ,NBW)
C
C      FORMULATE THE RIVER MATRIX TERMS
C
C      IF (LRIVER) CALL QRIVER (A,NIJ,NBW)
C      APPLY CONDENSATION OF SPECIFIED HEAD NODES TO MATRIX A
C      CALL BNDCON (A,UV,NIJ,NBW)
C      DECOMPOSE THE FULL MATRIX A TO UT*U
C      IF (.NOT.WTABLE) CALL SDCOMP (A,NIJ,NBW)
C      IF (WTABLE) CALL ADCOMP(A,NIJ,NBW)
C      FORMULATE THE WELL TERMS

```

```

MAI 260
MAI 270
MAI 280
MAI 290
MAI 300
MAI 310
MAI 320
MAI 330
MAI 340
MAI 350
MAI 360
MAI 370
MAI 380
MAI 390
MAI 400
MAI 410
MAI 420
MAI 430

```

```

C      CALL QZEE (UV)
C      CALL PUMPNU(UV)
C      IF(LPLOT) CALL PLOT(0,0,999)
C
C      BEGIN TIME STEP LOOP
C
C      DO 80 NN=1,NSTEP
C      WRITE(IO,4000)
C      4000 FORMAT (1H,////,5X,60('-',))
C      WRITE(IO,5000) NN
C      5000 FORMAT(1H0,4X,'TIME STEP NUMBER ',I6,/)
C      WRITE(IO,3000) MON(KMN),KDY,KYR
C      3000 FORMAT(1H,7X,'TIME STEP BEGINNING ',A3,I3,'',I5,/)
C
C      CALL QBOUND(A,UV,NIJ,NBW)
C      CALL LOADQR(UV)
C      TBH=0.0
C      DO 15 IM=1,12
C      15 TBH=TBH+PBH(IM)*FMON(IM)

```

```

MAI 440
MAI 451
MAI 460
MAI 470
MAI 480
MAI 490
JEC
JEC
JEC
JEC
MAI 500
MAI 510
MAI 520
MAI 530
MAI 531
MAI 532

```

```

C      INITIALIZE HI(I)
C
C      DO 20 I=1,NIJ
C      HI(I)=H(I)
C      20 CONTINUE
C
C      BEGIN ITERATIVE LOOP
C
C      DO 60 ITER=1,MAXIT
C      DO 30 L=1,NIJ
C      H(L)=HI(L)+BETA*(H(L)-HI(L))
C      HI(L)=H(L)
C      30 QZ(L)=QRE(L)
C      IF (LRIVER) CALL RIVER(A,NIJ,NSW)
C      DO 40 I=1,NIJ
C      IF (WTABLE) QZ(I)=QZ(I)-QBAR(I)
C      IF (IBNODE(I).LT.2) GO TO 40
C      QB(I)=QZ(I)
C      IF (WTABLE) QB(I)=QB(I)+QBAR(I)
C      QZ(I)=TBH*8QH(I)
C      IF (WTABLE) QZ(I)=QZ(I)-H0(I)
C      40 CONTINUE
C      IF (.NOT.WTABLE) CALL SSOLVE (A,NIJ,NSW,QZ,H)
C      IF (WTABLE) CALL ASOLVE (A,NIJ,NSW,QZ,H)
C
C      ERR=0.0
C      DO 50 L=1,NIJ
C      IF (WTABLE) H(L)=H(L)+H0(L)
C      HER=H(L)-HI(L)
C      HERR=HER*HER
C      IF (HERR.LE.ERR) GO TO 50
C      ERR=HERR
C      HMAX=HER
C      50 CONTINUE
C      IF (IPRINT.GE.2) WRITE (IO,6000) HMAX
C      HERR=EPS*EPS
C      IF (ERR.LT.HERR) GO TO 70
C      60 CONTINUE
C      70 WRITE(IO,8000)

```

```

MAI 540
MAI 550
MAI 560
MAI 570
MAI 580
MAI 590
MAI 600
MAI 610
MAI 620
MAI 630
MAI 640
MAI 650
MAI 660
MAI 670
MAI 680
MAI 690
MAI 700
MAI 710
MAI 720
MAI 730
MAI 740
MAI 750
MAI 760
MAI 770
MAI 780
MAI 790
MAI 800
MAI 810
MAI 820
MAI 830
MAI 840
MAI 850
MAI 860
MAI 870
MAI 880
MAI 890
MAI 900
MAI 910
MAI 920
MAI 930

```



```

C      READ (IN,5050) (PBH(I),I=1,12)
C      IF (IPRINT.GE.3) WRITE (IO,6110) (PBH(I),I=1,12)
C      READ (IN,5050) (PBF(I),I=1,12)
C      IF (IPRINT.GE.3) WRITE (IO,6111) (PBF(I),I=1,12)
C
C      GENERATE F.E. GRID. ONLY X,Y OR LAT, LONG OF THE ENDPOINTS OF EACH
C      VERTICAL LINE HAVE TO BE READ IN.
C
C      JT1=JT-1
C      AVGLAT = 0.0
C      DO 40 I=1,IS
C      NODBOT=(I-1)*JT+1
C      NODTOP=NODBOT+JT1
C      IF (LATLON) GO TO 20
C      READ (IN,6010) X(NODBOT),Y(NODBOT),X(NODTOP),Y(NODTOP)
C      X(NODBOT)=X(NODBOT)*XSCL
C      Y(NODBOT)=Y(NODBOT)*YSCL
C      X(NODTOP)=X(NODTOP)*XSCL
C      Y(NODTOP)=Y(NODTOP)*YSCL
C      GO TO 30
C      20 READ (IN,5010) LONB,LATB,LONT,LATT

```

C

```

C      PLONG=LONB(1)+(LONB(2)+(LONB(3))/60.)/60.
C      PLAT=LATB(1)+(LATB(2)+(LATB(3))/60.)/60.
C      AVGLAT = AVGLAT + PLAT/(2.0*IS)
C      CALL LAMBRT (PLONG,PLAT,XP,YP)
C      X(NODBOT)=XP
C      Y(NODBOT)=YP
C      PLONG=LONT(1)+(LONT(2)+(LONT(3))/60.)/60.
C      PLAT=LATT(1)+(LATT(2)+(LATT(3))/60.)/60.
C      AVGLAT = AVGLAT + PLAT/(2.0*IS)
C      CALL LAMBRT (PLONG,PLAT,XP,YP)
C      X(NODTOP)=XP
C      Y(NODTOP)=YP
C      30 IF (X(NODBOT).LT.X0) X0=X(NODBOT)
C      IF (X(NODTOP).LT.X0) X0=X(NODTOP)
C      IF (Y(NODBOT).LT.Y0) Y0=Y(NODBOT)
C      IF (Y(NODTOP).LT.Y0) Y0=Y(NODTOP)
C      DELX=(X(NODTOP)-X(NODBOT))/JT1
C      DELY=(Y(NODTOP)-Y(NODBOT))/JT1

```



```

        NGL(3,NE)=NO+JT+1
        GO TO 80
    60 NGL(1,NE)=NO
        NGL(2,NE)=NO+JT
        NGL(3,NE)=NO+JT+1
        NE=NE+1
        NGL(1,NE)=NO
        NGL(2,NE)=NO+JT+1
        NGL(3,NE)=NO+1
        GO TO 80
    70 NGL(1,NE)=NO
        NGL(2,NE)=NO+JT
        NGL(3,NE)=NO+JT+1
        NGL(4,NE)=NO+1
    80 CONTINUE

C      CALCULATE ELEMENT AREAS
C
C      SAREA=0.0
        IF (IPRINT.GE.5) WRITE (IO,6150)
        DO 100 NE=1,NEL
            N1=NGL(1,NE)
            N2=NGL(2,NE)
            N3=NGL(3,NE)
            IF (QUAD) GO TO 90
            AREA(NE)=0.5*((X(N2)-X(N1))*(Y(N3)-Y(N1))-(Y(N2)-Y(N1))*(X(N3)-X(N1)))
        11))
            IF (IPRINT.GE.5) WRITE (IO,6160) NE,N1,N2,N3,AREA(NE)
            GO TO 100
    90 N4=NGL(4,NE)
        AREA(NE)=0.5*((X(N1)-X(N3))*(Y(N2)-Y(N4))-(X(N2)-X(N4))*(Y(N1)-Y(N3)))
        13))
        IF (IPRINT.GE.5) WRITE (IO,6170) NE,N1,N2,N3,N4,AREA(NE)
    100 SAREA=SAREA+AREA(NE)
        IF (IPRINT.GE.2) WRITE (IO,6180) SAREA
C
C      CALL GRID PLOTTER

```

```

DIN1950
DIN1960
DIN1970
DIN1980
DIN1990
DIN2000
DIN2010
DIN2020
DIN2030
DIN2040
DIN2050
DIN2060
DIN2070
DIN2080
DIN2090
DIN2100
DIN2110
DIN2120
DIN2130
DIN2140
DIN2150
DIN2160
DIN2170
DIN2180
DIN2190
DIN2200
DIN2210
DIN2220
DIN2230
DIN2240
DIN2250
DIN2260
DIN2270
DIN2280
DIN2290
DIN2300
DIN2310

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```

C      IF (LPLOT) CALL PLTG
C      IF (.NOT.STEADY) GO TO 120
C      SET PAST VALUES OF HEAD ARRAY TO PRESENT VALUE FOR STEADY-STATE CADIN2320
C      DO 110 I=1,NIJ
C      110 HP(I)=H(I)
C      SAVE ORIGINAL HEAD VALUES FOR CALCULATING CUMULATIVE DRAWDOWN
C      120 IF (.NOT.DRAWD) RETURN
C      DO 130 I=1,NIJ
C      130 OLDH(I)=H(I)
C      RETURN
C      READ AND WRITE FORMATS
C      5000 FORMAT (10(L4,2X))
C      5010 FORMAT (6I4,6X,6I4)
C      5020 FORMAT (12A6,8X)
C      5030 FORMAT (20I4)
C      5040 FORMAT (I4,2F10.0,4I4)
C
C      6045 FORMAT (2F10.0)
C      6050 FORMAT (12F6.0)
C      6000 FORMAT (1H0,25X,12A6,/25X,12A6)
C      6003 FORMAT(1H1,/)
C      6005 FORMAT(1H,130('---'))
C      6010 FORMAT (8F10.0)
C      6020 FORMAT (1H0,30X,'LOGICAL CONTROL VARIABLES'/1H,5X,'STEADY',3X,'WT
DIN2550
DIN2560
DIN2570
DIN2580
DIN2590
DIN2600
DIN2610
DIN2620
DIN2630
1ABLE',5X,'QUAD',3X,'LATLON',3X,'L RIVER',4X,'DRAWD',4X,'LPLOT',3X,'
2EVPTRN'/1H,8X,8(L3,6X))
6030 FORMAT (1H0,5X,'THE PRINTOUT CONTROL PARAMETER (IPRINT) = ',I4/1H
1,5X,'NUMBER OF TIME STEPS BETWEEN EXTENDED PRINTOUTS (NPRNT)',' =
2',I4)
6035 FORMAT(1H0,5X,'NUMBER OF NODES IN ROW = ',I4,/6X,'NUMBER OF NODES
IN COLUMN = ',I4)

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6040 FORMAT (1H0,5X,'NUMBER OF NODES (NIJ) = ',I5/1H,5X,'NUMBER OF ', DIN2640
1'ELEMENTS (NEL) = ',I5/1H,5X,'HALF BANDWIDTH (NBW) = ',I3) DIN2650
6050 FORMAT (1H0,5X,'THE STARTING DATE FOR THE SIMULATION IS : ',A3,I3,DIN2660
1',',I5/1H,5X,'NUMBER OF DAYS PER TIME STEP (IDAY) = ',I2/1H,5X,'DIN2670
2NUMBER OF TIME STEPS TO BE TAKEN (NSTEP) = ',I4) DIN2680
6060 FORMAT (1H0,30X,'ITERATION PARAMETERS FOR STREAMFLOW SOLUTION',',', RDIN2690
1OUTINE',/1H,5X,'MAXIMUM NUMBER OF ITERATIONS (MAXIT) = ',I3/1H,5X,DIN2700
2,'CONVERGENCE ACCELERATION PARAMETER (BETA) = ',F8.5/1H,5X,'CONVEDIN2710
3RGENCE CRITERION (EPS) = ',F8.6) DIN2720
6070 FORMAT (1H0,5X,'SCALING FACTOR FOR X - COORDINATES (XSCL) = ',F12.DIN2730
13/1H,5X,'SCALING FACTOR FOR Y - COORDINATES (YSCL) = ',F12.3) DIN2740
6080 FORMAT (1H0,5X,'PLOT CONTROL PARAMETER (IPLOT) = ',I2/1H,5X,'PLOTDIN2750
1 SCALING FACTOR IN X DIRECTION (XSCLP) = ',F10.3/1H,5X,'PLOT SCALDIN2760
2ING FACTOR IN Y DIRECTION (YSCLP) = ',F10.3) DIN2770
6090 FORMAT (1H1,30X,'BOUNDARY-CONDITION TYPE MAP',/1H,35X,'0 - INTERIODIN2780
1R NODE',/1H,35X,'1 - SPECIFIED-FLUX BOUNDARY NODE',/1H,35X,'2 - SPDIN2790
2ECIFIED-HEAD BOUNDARY NODE',/) DIN2800
6100 FORMAT(1H,2X,I4,2X,I4,4X,(20I4)) DIN2820
6110 FORMAT (1H0,30X,'MONTHLY MULTIPLIERS OF SPECIFIED HEAD VALUES',/ DIN2830
11H,5X,12F8.2) DIN2831
6111 FORMAT (1H0,30X,'MONTHLY MULTIPLIERS OF SPECIFIED FLUX VALUES',/ DIN2832
11H,5X,12F8.2) DIN2840
6120 FORMAT (1H1,10X,'THE X AND Y OFFSETS FOR THE GRID ORIGIN ARE ',2F1DIN2840
15.4) DIN2850
6130 FORMAT (1H0,30X,'NODAL X - Y COORDINATES',/) DIN2860
6140 FORMAT (1H,5X,I4,2F12.4) DIN2870
6150 FORMAT (1H1,30X,'GLOBAL INDICES AND ELEMENT AREAS',/) DIN2880
6160 FORMAT (1H,5X,4I4,F15.4) DIN2890
6170 FORMAT (1H,5X,5I4,F15.4) DIN2900
6180 FORMAT (1H0,10X,'TOTAL MODEL AREA = ',1P1E15.5) DIN2910
6190 FORMAT(1H,2X,' ROW',2X,'NODE')
END
DIN2920-

```


C

```

      IF (MOD(KYR,4).EQ.0) IDMON(2)=29
      GO TO 70
80  KD=KDY+IDAY
      ID=IDAY
      KM=KMN
      P1=1.0
90  M2=0
      IF (KD.GT.IDMON(KM)) M2=KD-IDMON(KM)
      M1=ID-M2
      FMON(KM)=(1.0*M1)/(1.0*IDAY)
      KM1=KM+1
      IF (KM1.GT.12) KM1=KM1-12
      FMON(KM1)=P1-FMON(KM)
      IF (M2.LE.IDMON(KM)) GO TO 100
      ID=ID-M1
      P1=P1-FMON(KM)
      KM=KM1
      KD=M2
      GO TO 90
100 RETURN
      END

```

```

DTE 660
DTE 670
DTE 680
DTE 690
DTE 700
DTE 710
DTE 720
DTE 730
DTE 740
DTE 750
DTE 760
DTE 770
DTE 780
DTE 790
DTE 800
DTE 810
DTE 820
DTE 830
DTE 840
DTE 850
DTE 860-

```



```

C
WRITE (IO,6030) INAME
NODE=0
IF (NUNIT.LT.1) GO TO 40
C
READ IN ARRAY OF NODAL VALUES
C
WRITE (IO,6000) TITLE,NUNIT,FMT,SCALE
WRITE(IO,6035)
DO 30 IR=1,IS
READ (NUNIT,FMT) (X(NODE+L),L=1,JT)
IF (SCALE.EQ.0..OR.SCALE.EQ.1.) GO TO 20
DO 10 L=1,JT
NODEL=NODE+L
10 X(NODEL)=X(NODEL)*SCALE
20 NNODE = NODE + 1
IF(IPRNT.GT.0) WRITE(IO,6010) IR,NNODE,(X(NODE+L),L=1,JT)
NODE=NODE+JT
30 CONTINUE
RETURN
40 CONTINUE
C
SET ALL NODAL VALUES TO 'SCALE', AND READ NO OTHER CARDS
C
DO 50 L=1,NIJ
50 X(L)=SCALE
WRITE (IO,6020) TITLE,SCALE
RETURN
C
5000 FORMAT (2X,1A6,3X,I2,2X,2A6,2X,F10.0,I4)
6000 FORMAT (1H0,2X,1A6,20H WAS READ FROM UNIT ,I2,11H,ON FORMAT ,2A6,1
16H, AND SCALED BY ,1P1E12.3/)
6010 FORMAT(1H ,2X,I4,2X,I4,2X,(1P10E10.3))
6020 FORMAT (1H0,2X,1A6,42H WAS NOT READ, ENTIRE MATRIX SET EQUAL TO ,1
1P1E12.3)
6030 FORMAT(1H1,30X,'INPUT DATA FOR ARRAY ',A6)
6035 FORMAT(1H ,2X,' ROW',2X,'NODE')
END
C
ARY 350
ARY 360
ARY 370
ARY 380
ARY 390
ARY 400
ARY 410
ARY 420
ARY 430
ARY 440
ARY 450
ARY 460
ARY 470
ARY 490
ARY 500
ARY 510
ARY 520
ARY 530
ARY 540
ARY 550
ARY 560
ARY 570
ARY 580
ARY 590
ARY 600
ARY 610
ARY 620
ARY 630
ARY 650
ARY 660
ARY 680-

```



```

C
20  XP=X(I*JT)/XSCLP
   YP=Y(I*JT)/YSCLP
   CALL PLOT (XP,YP,3)
   XP=X((I-1)*JT+1)/XSCLP
   YP=Y((I-1)*JT+1)/YSCLP
   CALL PLOT (XP,YP,2)
30  CONTINUE
C
   DRAW BOTTOM LINE
C
   ISM=ISM-1
   IP=ISM*JT+1
   XP=X(IP)/XSCLP
   YP=Y(IP)/YSCLP
   CALL PLOT (XP,YP,3)
   DO 40 J=1,ISM
   IP=IP-JT
   XP=X(IP)/XSCLP
   YP=Y(IP)/YSCLP
   CALL PLOT (XP,YP,2)
40  CONTINUE
C
   DRAW HORIZONTAL LINES
C
   DO 70 I=2,JT
   XP=X(I)/XSCLP
   YP=Y(I)/YSCLP
   CALL PLOT (XP,YP,3)
   DO 50 J=1,ISM
   IP=J*JT+I
   XP=X(IP)/XSCLP
   YP=Y(IP)/YSCLP
   CALL PLOT (XP,YP,2)
50  CONTINUE
   IF (QUAD) GO TO 70
C

```

```

PLT 410
PLT 420
PLT 430
PLT 440
PLT 450
PLT 460
PLT 470
PLT 480
PLT 490
PLT 500
PLT 510
PLT 520
PLT 530
PLT 540
PLT 550
PLT 560
PLT 570
PLT 580
PLT 590
PLT 600
PLT 610
PLT 620
PLT 630
PLT 640
PLT 650
PLT 660
PLT 670
PLT 680
PLT 690
PLT 700
PLT 710
PLT 720
PLT 730
PLT 740
PLT 750
PLT 760

```

PLT 770
PLT 780
PLT 790
PLT 800
PLT 810
PLT 820
PLT 830
PLT 840
PLT 850
PLT 860
PLT 870
PLT 880
PLT 890
PLT 900
PLT 910
PLT 920
PLT 930
PLT 940
PLT 950
PLT 960
PLT 970
PLT 980
PLT 990
PLT1000

DRAW LINES FOR TRIANGULAR ELEMENTS

```

      IZ=JT*ISM+I
      IF (MOD(IS,2).EQ.0) IZ=IZ-1
      XP=X(IZ)/XSCLP
      YP=Y(IZ)/YSCLP
      CALL PLOT (XP,YP,3)
      DO 60 J=1,ISM
      ISP=IS-J
      KP=JT-1
      IF (MOD(ISP,2).EQ.0) KP=JT+1
      IZ=IZ-KP
      XP=X(IZ)/XSCLP
      YP=Y(IZ)/YSCLP
      CALL PLOT (XP,YP,2)
60    CONTINUE
70    CONTINUE
      IF (IPLOT.LT.4) RETURN

```

PUT ELEMENT NO. AT MIDPOINT

```

      DO 90 NE=1,NEL
      XBAR=0.0
      YBAR=0.0

```

PLT1010
PLT1020
PLT1030
PLT1040
PLT1050
PLT1060
PLT1070
PLT1080
PLT1090
PLT1100

```

      DO 80 J=1,NETYP
      N1=NGL(J,NE)
      XBAR=XBAR+X(N1)
      YBAR=YBAR+Y(N1)
80    CONTINUE
      YBAR=YBAR/(1.0*NETYP*YSCLP)
      XBAR=XBAR/(1.0*NETYP*XSCLP)
      FNN=NE*1.
      CALL NUMBER (XBAR,YBAR,.07,FNN,0.,-1)
90    CONTINUE
100   RETURN
      END

```

PLT1120
PLT1130


```

DO 50 IP=1,4
ETA=E(IP)
ZETA=Z(IP)
DJ=AREA2+ETA*(X34*Y12-X12*Y34)+ZETA*(X23*Y14-X14*Y23)
DJAC=0.125*DJ
DJ=1./DJ
F1 = 1.-ETA
F2 = 1.+ETA
F3 = 1.-ZETA
F4 = 1.+ZETA
W(1) = .25*F1*F3
W(2) = .25*F2*F3
W(3) = .25*F2*F4
W(4) = .25*F1*F4
W(1)=0.25
W(2)=0.25
W(3)=0.25
W(4)=0.25
WY(1)=(-X24+X34*ETA+X23*ZETA)*DJ
WY(2)=(X13-X34*ETA-X14*ZETA)*DJ
WY(3)=(X24-X12*ETA+X14*ZETA)*DJ
WY(4)=(-X13+X12*ETA-X23*ZETA)*DJ
WX(1)=(Y24-Y34*ETA-Y23*ZETA)*DJ
WX(2)=(-Y13+Y34*ETA+Y14*ZETA)*DJ
WX(3)=(-Y24+Y12*ETA-Y14*ZETA)*DJ
WX(4)=(Y13-Y12*ETA+Y23*ZETA)*DJ
TRAN=0.0
GRADX=0.0
GRADY=0.0
STOR=0.0
DO 30 I=1,NETYP
NI=N(I)
WI=W(I)
IF (.NOT.STEADY) STOR=STOR+S(NI)*WI
TR=T(NI)*WI
IF (.NOT.WTABLE) GO TO 30
TR=TR*(HO(NI)-BOT(T(NI)))
GRADX=GRADX+HO(NI)*WX(I)
GRADY=GRADY+HO(NI)*WY(I)

```

LMQ 450
LMQ 460
LMQ 470
LMQ 480
LMQ 490
LMQ 500
LMQ 510
LMQ 520
LMQ 530
LMQ 540
LMQ 550
LMQ 560
LMQ 570
LMQ 580
LMQ 590
LMQ 600
LMQ 610
LMQ 620
LMQ 630
LMQ 640
LMQ 650
LMQ 660
LMQ 670
LMQ 680
LMQ 690
LMQ 700
LMQ 710
LMQ 720
LMQ 730
LMQ 740
LMQ 750
LMQ 760
LMQ 770
LMQ 780
LMQ 790
LMQ 800
LMQ 810
LMQ 820
LMQ 830

```

30  TRAN=TRAN+TR
    TRAN=TRAN*DJAC
    STOR=STOR*DJAC
    DO 50 I=1,NETYP
      NI=N(I)
      WI=W(I)
      WXI=WX(I)
      WYI=WY(I)
      IF (.NOT. STEADY) UV(NI)=UV(NI)+STOR*WI
      DO 50 J=1,NETYP
        NJ=N(J)
        IF (WTABLE) GO TO 40
        IF (NJ.LT.NI) GO TO 50
        NJB=NJ-NI+IDIAG
        A(NI,NJB)=A(NI,NJB)+TRAN*(WXI*WX(J)+WYI*WY(J))
        GO TO 50
40   WJ=W(J)
      WXJ=WX(J)
      WYJ=WY(J)
      TR=TRAN*(WXI*WXJ+WYI*WYJ)
      NJE=NJ-NI+IDIAG

```

C

```

    A(NI,NJB)=A(NI,NJB)+TR+T(NJ)*WJ*(GRADX*WXI+GRADY*WYI)*DJAC
    QBAR(NI)=QBAR(NI)+TR*HO(NJ)
50  CONTINUE
60  CONTINUE
    IF (STEADY) RETURN
    RDT=0.5/DELT
    DO 70 IJ=1,NIJ
      UV(IJ)=RDT*UV(IJ)
70  A(IJ,IDIAG)=A(IJ,IDIAG)+3.*UV(IJ)
    RETURN
    END

```

LMQ 840
 LMQ 850
 LMQ 860
 LMQ 870
 LMQ 880
 LMQ 890
 LMQ 900
 LMQ 910
 LMQ 920
 LMQ 930
 LMQ 940
 LMQ 950
 LMQ 960
 LMQ 970
 LMQ 980
 LMQ 990
 LMQ1000
 LMQ1010
 LMQ1020
 LMQ1030
 LMQ1040

LMQ1050
 LMQ1060
 LMQ1070
 LMQ1080
 LMQ1090
 LMQ1100
 LMQ1110
 LMQ1120
 LMQ1130
 LMQ1140
 LMQ1150-

LMT 380
LMT 390
LMT 400
LMT 410
LMT 420
LMT 430
LMT 440
LMT 450
LMT 460
LMT 470
LMT 480

DO 60 NE=1,NEL
N1=NGL(1,NE)
N2=NGL(2,NE)
N3=NGL(3,NE)
N(1)=N1
N(2)=N2
N(3)=N3
X23=X(N2)-X(N3)
Y23=Y(N3)-Y(N2)
X31=X(N3)-X(N1)

C

C

LMT 490
LMT 500
LMT 510
LMT 520
LMT 530
LMT 540
LMT 550
LMT 560
LMT 570
LMT 580
LMT 590
LMT 600
LMT 610
LMT 620
LMT 630

Y31=Y(N1)-Y(N3)
X12=X(N1)-X(N2)
Y12=Y(N2)-Y(N1)
DJAC=AREA(NE)
DJ=0.5/DJAC
W(1)=1.0/3.0
W(2)=1.0/3.0
W(3)=1.0/3.0
WY(1)=X23*DJ
WY(2)=X31*DJ
WY(3)=X12*DJ
WX(1)=Y23*DJ
WX(2)=Y31*DJ
WX(3)=Y12*DJ

C

C

LMT 640
LMT 650
LMT 660
LMT 670
LMT 680

TRAN - MEAN TRANSMISSIVITY OVER AN ELEMENT (FOR A WATER-TABLE
AQUIFER, TRAN IS AVG. PERMEABILITY TIMES THE THICKNESS BASED ON
STOR - MEAN STORAGE TERM OVER AN ELEMENT
GRADX - MEAN GRADIENT IN X OF 'HO' OVER AN ELEMENT
GRADY - MEAN GRADIENT IN Y OF 'HO' OVER AN ELEMENT

C

C

C

C

C

C

C

C

```

TRAN=0.0
GRADX=0.0
GRADY=0.0
STOR=0.0
DO 30 I=1,NETYP
  NI=N(I)
  WI=W(I)
  IF (.NOT.STEADY) STOR=STOR+S(NI)*WI
  TR=T(NI)*WI
  IF (.NOT.WTABLE) GO TO 30
  TR=TR*(HO(NI)-BOT(T(NI)))
  GRADX=GRADX+HO(NI)*WX(I)
  GRADY=GRADY+HO(NI)*WY(I)
30  TRAN=TRAN+TR
    TRAN=TRAN*DJAC
    STOR=STOR*DJAC

C
C   LOAD LOCAL ELEMENT TERMS INTO GLOBAL COEFFICIENT MATRICES. FOR
C   CONFINED AQUIFER, LOAD ONLY THE SYMMETRIC HALFBAND.
C
DO 50 I=1,NETYP
  NI=N(I)
  WI=W(I)
  WXI=WX(I)
  WYI=WY(I)
  IF (.NOT.STEADY) UV(NI)=UV(NI)+STOR*WI
DO 50 J=1,NETYP
  NJ=N(J)
  IF (WTABLE) GO TO 40
  IF (NJ.LT.NI) GO TO 50
  NJB=NJ-NI+IDIAG
  A(NI,NJB)=A(NI,NJB)+TRAN*(WXI*WX(J)+WYI*WY(J))
GO TO 50
40  WXJ=WX(J)
    WYJ=WY(J)
    TR=TRAN*(WXI*WXJ+WYI*WYJ)
    NJB=NJ-NI+IDIAG

```

C


```

A(NI,NJB)=A(NI,NJB)+TR+T(NJ)*(GRADX*WXI+GRADY*WYI)*DJAC
QBAR(NI)=QBAR(NI)+TR*HO(NJ)
50 CONTINUE
60 CONTINUE
IF (STEADY) RETURN

FOR TRANSIENT CASE, LOAD TIME VECTOR INTO GLOBAL 'A' MATRIX

RDT=0.5/DELT
DO 70 I=1,NIJ
UV(I)=UV(I)*RDT
A(I,IDIAG)=A(I,IDIAG)+3.0*UV(I)
70 CONTINUE
RETURN
END

-----
SUBROUTINE QRIVER (A,NIJD,NBWD)
-----
IMPLICIT DOUBLE PRECISION (A-H,O-Z)
REAL BUDTTL,LENGTH
LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON/AQUIFR/ H(572),HP(572),QRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
COMMON/WTBLE/ QBAR(572),BOTT(572),HO(572)
COMMON /BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
COMMON/RIVERB/ HPOOL(2),STAGE(20,2),DSCHRG(20,2),QCOMR(118),
1ELVR(118),XFACR(118),RW(4,118),RIVERH(118),
2RIVERQ(118),FLOWIN(366,7),QCOR(5),SEWAGE,NODEIN(7),NSTAGE(2),
3NRN(118),NER(118),NRTYP(118),NREACH,NRNODE,NTRIB,QRR(64)
COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYP
COMMON /MPSC/ XSCL,YSCL,XO,YO
COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KDYS,IDYR,MON(12)
COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON /IUNIT/ IN,IO,IPRINT,NPRNT
DIMENSION A(NIJD,NBWD),LON(3),LAT(3)

```

LMT1070
LMT1080
LMT1090
LMT1100
LMT1110
LMT1120
LMT1130
LMT1140
LMT1150
LMT1160
LMT1170
LMT1180
LMT1190
LMT1200
LMT1210-

QRV 10

QRV 20

QRV 30
QRV 60
QRV 70
QRV 80
QRV 90
QRV 100
QRV 110
QRV 120*

QRV 140
QRV 150
MAI 140
QRV 170
QRV 180
QRV 190
QRV 200

```

C      RIVER CONTROL INTEGERS
C      -----
C      NREACH - THE NUMBER OF REACHES THE RIVER IS DIVIDED INTO
C      NRNODE - TOTAL NO. OF RIVER NODES
C      NRZ - THE NUMBER OF ZONES THE RIVER IS DIVIDED INTO BASED ON THE
C      THE VALUE FOR LEAKANCE. THE ZONES DO NOT HAVE TO BE CONT
C      QCOR - LEAKANCE VALUE FOR A ZONE ( K'/B' )
C      NSTAGE -
C      HPOOL - THE POOL ELEVATION OF THE RESEVIOR (ONE PER REACH)
C      STAGE,DSCHRG - POINTS DESCRIBING THE LINEAR APPROXIMATION OF THE
C      STAGE-DISCHARGE RELATIONSHIP FOR THE REACH
C      -----
C      READ (IN,5000) NREACH,NRNODE,NRZ
C      READ (IN,5000) NREACH,NRNODE,NRZ,NTRIB
C      IF (IPRINT.GE.2) WRITE (IO,6000) NREACH,NRNODE,NRZ
C      NBUD=12+NREACH
C      IF (IPRINT.GE.2) WRITE (IO,6010)
C      READ (IN,5025) (QCOR(I),I=1,NRZ)
C      IF (IPRINT.GE.2) WRITE (IO,6020) (I,QCOR(I),I=1,NRZ)
C      IF (IPRINT.GE.2) WRITE (IO,6030)
C      DO 20 N=1,NREACH
C      READ (IN,5010) NSTAGE(N),HPOOL(N)
C      IF (IPRINT.GE.2) WRITE (IO,6040) N,NSTAGE(N),HPOOL(N)
C
C
C      NSTAGN=NSTAGE(N)
C      READ (IN,5020) (STAGE(L,N),L=1,NSTAGN)
C      READ (IN,5020) (DSCHRG(L,N),L=1,NSTAGN)
C      IF (IPRINT.GE.2) WRITE (IO,6050) (STAGE(L,N),L=1,NSTAGN)
C      20 IF (IPRINT.GE.2) WRITE (IO,6060) (DSCHRG(L,N),L=1,NSTAGN)
C      NR=0
C      IF (IPRINT.GE.4) WRITE (IO,6070)
C      DO 60 NODE=1,NRNODE
C      DO 50 NODE=1,NRNODE
C      IF (LATLON) GO TO 30
C      READ (IN,5040) XR,YR,NRN(NODE),IRZ,NER(NODE),NRTP(NODE),LENGTH,WID
C      1TH,ELVR(NODE),XFACTR(NODE)
C      XR=XR*XSCL

```

```

C      YR=YR*YSCL
      GO TO 40
30 READ (IN,5030) LON,LAT,NRN(NODE),IRZ,NER(NODE),NRTYP(NODE),LENGTH,W
      1IDTH,ELVR(NODE),XFACTR(NODE)
      PLONG=LON(1)+(LON(3)/60.+LON(2))/60.
      PLAT=LAT(1)+(LAT(3)/60.+LAT(2))/60.
      CALL LAMBERT (PLONG,PLAT,XR,YR)
40 QCOMR(NODE)=QCOR(IRZ)*LENGTH*WIDTH
      XR = XSCL*(XR-XO)
      YR = YSCL*(YR-YO)
      IF (LPLOT) CALL PLTR (XR,YR,NODE)
      NERN=NER(NODE)
      IF (IPRINT.GE.4) WRITE (IO,6080) NODE,XR,YR,NRN(NODE),IRZ,NERN,NRTY
      1P(NODE),LENGTH,WIDTH,ELVR(NODE),XFACTR(NODE)
      CALL PTFNDR (XR,YR,NERN,RW,NODE,1)

C      QCOMRN=QCOMR(NODE)
      DO 50 I=1,NETYP
      NI=NGL(I,NERN)
      RWI=RW(I,NODE)
      DO 50 J=1,NETYP
      NJ=NGL(J,NERN)
      IF (.NOT.WTABLE.AND.NJ.LT.NI) GO TO 50
      NJB=NJ-NI+IDIAG
      A(NI,NJB)=A(NI,NJB)+QCOMRN*RWI*RW(J,NODE)
      IF (WTABLE) QBAR(NI)=QBAR(NI)+QCOMRN*RWI*RW(J,NODE)*MO(NJ)
50 CONTINUE
C      IF (NRN(NODE).LT.(NR+2)) GO TO 60
      NR=NR+1
      NEND(NR)=NODE
C      60 CONTINUE
      NEND(NR+1)=NRNODE+1
      IF (NTRIB.LT.1) RETURN
      READ (IN,5000) (NODEIN(N),N=1,NTRIB)
      WRITE(IO,5005) (NODEIN(N),N=1,NTRIB)
      DO 60 N=1,NTRIB
C
C      READ DAILY INFLOW
      READ (IN,5015) (FLOWIN(L,N),L=1,366)
      IF (IPRINT.GE.2) WRITE (IO,6090) N
C      60 IF (IPRINT.GE.2) WRITE (IO,6100) (FLOWIN(L,N),L=1,366)

```

```

C      READ (IN,5015) (FLOWIN(N),N=1,366)
C      IF (IPRINT.GE.2) WRITE (IO,6090)
C      IF (IPRINT.GE.2) WRITE (IO,6100) (FLOWIN(N),N=1,366)
5000  FORMAT (20I4)
5005  FORMAT(1H0,5X,'RIVER-NODE LOCATION OF TRIBUTARIES ',20I4,/)

```

```

QRV 870
QRV 880
QRV 890
QRV1490
JEC1495

```

C

```

5010  FORMAT (I4,F10.0)
5015  FORMAT (12F10.0)
5020  FORMAT (12F6.0)
5025  FORMAT (7F10.0)
5030  FORMAT (10I4,4F10.0)
5040  FORMAT (2F12.0,4I4,4F10.0)
6000  FORMAT (1H1/1H ,30X,'RIVER PARAMETERS',/1H0,5X,'NUMBER OF ',/,'RIVER
1REACHES = ',I3/1H ,5X,'NUMBER OF RIVER NODES = ',I4/1H ,5X,'NUMBER
2 OF RIVER-LEAKANCE ZONES = ',I3)
6010  FORMAT (1H0,30X,'RIVER-LEAKANCE VALUES BY ZONE (QCOR(N))')
6020  FORMAT (1H ,5X,'ZONE ',I2,2X,1P1E12.5)
6030  FORMAT (1H0,30X,'RIVER-REACH PARAMETERS')
6040  FORMAT (1H ,10X,'REACH NUMBER ',I2/1H ,5X,I3,' POINTS USED IN',/,'
1TAG - DISCHARGE RELATIONSHIP',/1H ,5X,'POOL ELEVATION = ',F10.3)
6050  FORMAT (1H ,5X,' STAGE ',12F9.3)
6060  FORMAT (1H ,5X,'DISCHARGE',12F9.3)
6070  FORMAT (1H0,30X,'RIVER-NODE DATA',/1H0,5X,'NODE',4X,'X-COORD',4X,'Y
1-COORD',3X,'REACH',2X,'ZONE',2X,'ELEM.',2X,'TYPE',5X,'LENGTH',5X,'
2WIDTH',5X,'ELEVATION',5X,'XFACTOR')
6080  FORMAT (1H ,5X,I3,2F12.3,4I6,5X,4F10.3)
C6090  FORMAT (1H0,30X,'DAILY INFLOW VALUES (FLOWIN(N))',/)
6090  FORMAT (1H0,30X,'DAILY INFLOW VALUES (FLOWIN(N))',3X,I2)
6100  FORMAT (1H ,12F10.2)
      RETURN
      END

```

```

QRV1500

QRV1520
QRV1530
QRV1540
QRV1550
QRV1560
QRV1570
QRV1580
QRV1590
SQRV1600
QRV1610
QRV1620
QRV1630
YQRV1640
QRV1650
QRV1660
      QRV1670
QRV1680
QRV1680*
QRV1690
QRV 900

```

```

C -----
SUBROUTINE RIVER (A,NIJD,NBWD)
C -----
IMPLICIT DOUBLE PRECISION (A-H,O-Z)
REAL BUDTTL,LENGTH
LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON/AQUIFR/ H(572),HP(572),QRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
COMMON/WTBLE/ QBAR(572),BOTT(572),H0(572)
COMMON/BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
COMMON/RIVERB/ HPOOL(2),STAGE(20,2),DSCHRG(20,2),QCOMR(118),
1ELVR(118),XFACR(118),RW(4,118),RIVERH(118),
2RIVERQ(118),FLOWIN(366,7),QCOR(5),SEWAGE,NODEIN(7),NSTAGE(2),
3NRN(118),NER(118),NRTYP(118),NREACH,NRNODE,NTRIS,QRR(64)
COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYP
COMMON /MPSC/ XSCL,YSCL,XO,YO
COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KDYS,IDYR,MON(12)
COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON /IOUNIT/ IN,IO,IPRINT,NPRNT
DIMENSION A(NIJD,NBWD),LON(3),LAT(3)
C ZERO RIVER FLOW ARRAY
DO 65 I=1,NRNODE
65 RIVERQ(I) = 0.0
C ZERO REACH BUDGET ARRAY
DO 66 L=1,NREACH
66 BUD(12+L) = 0.0
IF (NTRIS.LT.1) GO TO 77
DO 75 N=1,NTRIS
I = NODEIN(N)
RQ1=0.
JDAY=KDYS-1
DO 70 L=1,IDAY
JDAY=JDAY+1
IF (JDAY.GT.IDYR) JDAY=JDAY-IDYR
C 70 RQ1=RQ1+FLOWIN(JDAY)

```



```

150 Q=(RIVERH(NODE)+B-HAQ)*QCOMR(NODE)
    IF (Q.GT.RIVERQ(NODE)) Q=RIVERQ(NODE)
    QRR(NODE)=Q
    IF (NODE.GE.NRNODE) GO TO 155
    RIVERQ(NODE+1)=RIVERQ(NODE)-Q
    RIVERQ(NODE+1)=RIVERQ(NODE+1)+RIVERQ(NODE)-Q
    IF (RIVERQ(NODE+1).GE.0.0) GO TO 155
    WRITE (IO,6110) NODE,RIVERQ(NODE+1)
    RIVERQ(NODE+1) = 0.0
155 BUD(12+L) = BUD(12+L) + Q
    Q=Q+QCOMR(NODE)*HAQ
    DO 160 I=1,NETYP
    NI=NGL(I,NERN)
160 QZ(NI)=QZ(NI)+Q*RW(I,NODE)
170 CONTINUE
    BUD(13)=RIVERQ(1)-RIVERQ(NEND(1))
C
C
C IF (NREACH.LT.2) RETURN
C NR1=NREACH-1
C DO 180 N=1,NR1
C L=N+13
C 180 BUD(L)=RIVERQ(NEND(N))-RIVERQ(NEND(N+1))
6110 FORMAT (5X,'DIVERSION AFTER NODE',I3,'EXCEEDSFLOWBY',F10.2)
    RETURN
C

```

END

QRV1700-

QRV1320
QRV1330
JEC
QRV1335*
QRV1340
QRV1340*
QRV1341*
QRV1342*
QRV1343*
QRV1344*
QRV1350
QRV1360
QRV1370
QRV1380
QRV1390
QRV1400
QRV1410
QRV1420
QRV1430
QRV1440
QRV1450
QRV1460
QRV1691*
QRV1470


```

30  IF (D) 100,30,100
40  IF (V) 50,40,50
    4C  ETA=(G*X0-C*Y0)/(B*G-F*C)
    ZETA=(B*Y0-F*X0)/(B*G-F*C)
    GO TO 230
50  IF (C) 70,60,70
    60  ETA=X0/B
    ZETA=(Y0-F*ETA)/(G+V*ETA)
    GO TO 230
70  IF (B) 90,80,90
    80  ZETA=X0/C
    ETA=(Y0-G*ZETA)/(F+V*ZETA)
    GO TO 230
90  ABAR=-B*V/C
    BBAR=F+V*X0/C-B*G/C
    CBAR=G*X0/C-Y0
    BBAR=BBAR/(2.*ABAR)
    CBAR=CBAR/ABAR
    YBB=BBAR*BBAR-CBAR
    CSTAR=DSQRT(YBB)
    ETA=-BBAR+CBAR
    ZETA=(X0-B*ETA)/C
    R1=ETA*ETA+ZETA*ZETA
    IF (ETA*ETA.GT.1..OR.ZETA*ZETA.GT.1.) ETA=-BBAR-CSTAR
    ZETA=(X0-B*ETA)/C
    R2=ETA*ETA+ZETA*ZETA
    IF (R2.GT.R1) ETA=-BBAR+CBAR
    ZETA=(X0-B*ETA)/C
    GO TO 230
100 IF (V) 160,110,160
110 IF (G) 130,120,130
120 ETA=Y0/F
    ZETA=(X0-B*ETA)/(C+D*ETA)
    GO TO 230
130 IF (F) 150,140,150
140 ZETA=Y0/G
    ETA=(X0-C*ZETA)/(B+D*ZETA)

```

```

PTF 780
PTF 790
PTF 800
PTF 810
PTF 820
PTF 830
PTF 840
PTF 850
PTF 860
PTF 870
PTF 880
PTF 890
PTF 900
PTF 910
PTF 920
PTF 930
PTF 940
PTF 950
PTF 960
PTF 970
PTF 980
PTF 990
PTF1000
PTF1010
PTF1020
PTF1030
PTF1040
PTF1050
PTF1060
PTF1070
PTF1080
PTF1090
PTF1100
PTF1110
PTF1120
PTF1130
PTF1140
PTF1150
PTF1160

```

```

GO TO 230
150 ABAR=-D*F/G
BBAR=B+D*Y0/G-C*F/G
CBAR=C*Y0/G-X0
BBAR=BBAR/(2.*ABAR)
CBAR=CBAR/ABAR
YBB=BBAR*BBAR-CBAR
CSTAR=DSQRT(YBB)
ETA=-BBAR+CSTAR
ZETA=(Y0-F*ETA)/G
R1=ETA*ETA+ZETA*ZETA
IF (ETA*ETA.GT.1..OR.ZETA*ZETA.GT.1.) ETA=-BBAR-CSTAR
ZETA=(Y0-F*ETA)/G
R2=ETA*ETA+ZETA*ZETA
IF (R2.GT.R1) ETA=-BBAR+CSTAR
ZETA=(Y0-F*ETA)/G
GO TO 230
160 ABAR=B*V-F*D
BBAR=C*V-D*G
CBAR=V*X0-D*Y0
IF (BBAR) 200,170,200
170 IF (ABAR) 190,180,190
180 GO TO 230
190 ETA=CBAR/ABAR
ZETA=(Y0-F*ETA)/(V*ETA+G)
GO TO 230
200 IF (ABAR) 220,210,220
210 ZETA=CBAR/BBAR
ETA=(X0-C*ZETA)/(B+D*ZETA)
GO TO 230
220 XBAR=-ABAR*D/BBAR
YBAR=B+CBAR*D/BBAR-ABAR*C/BBAR
ZBAR=CBAR*C/BBAR-X0
YBAR=YBAR/(2.*XBAR)
ZBAR=ZBAR/XBAR
YBB=(YBAR*YBAR-ZBAR)
PTF1170
PTF1180
PTF1190
PTF1200
PTF1210
PTF1220
PTF1230
PTF1240
PTF1250
PTF1260
PTF1270
PTF1280
PTF1290
PTF1300
PTF1310
PTF1320
PTF1330
PTF1340
PTF1350
PTF1360
PTF1370
PTF1380
PTF1390
PTF1400
PTF1410
PTF1420
PTF1430
PTF1440
PTF1450
PTF1460
PTF1470
PTF1480
PTF1490
PTF1500
PTF1510
PTF1520

```

```

IF (YBB.LT.O.) YBB=O.O
CSTAR=DSQRT(YBB)
ETA=-YBAR+CSTAR
ZETA=(CBAR-ABAR*ETA)/BBAR
R1=ETA*ETA+ZETA*ZETA
IF (ETA*ETA.GT.1..OR.ZETA*ZETA.GT.1.) ETA=-YBAR-CSTAR
ZETA=(CBAR-ABAR*ETA)/BBAR
R2=ETA*ETA+ZETA*ZETA
IF (R2.GT.R1) ETA=-YBAR+CSTAR
ZETA=(CBAR-ABAR*ETA)/BBAR
230 CONTINUE
W(IP+1)=(1.-ETA)*(1.-ZETA)*.25
W(IP+2)=(1.+ETA)*(1.-ZETA)*.25
W(IP+3)=(1.+ETA)*(1.+ZETA)*.25
W(IP+4)=(1.-ETA)*(1.+ZETA)*.25
IF (ETA*ETA.LE.1.O.AND.ZETA*ZETA.LE.1.O) RETURN
C
IF THE MAGNITUDE OF ETA OR ZETA IS > 1, IT MUST BE LOCATED IN
A DIFFERENT ELEMENT
C
IF (NT.EQ.1) WRITE (IO,6020) NO,NE,ETA,ZETA
IF (NT.EQ.2) WRITE (IO,6030) NO,NE,ETA,ZETA
RETURN
C

```

```

PTF11530
PTF11540
PTF11550
PTF11560
PTF11570
PTF11580
PTF11590
PTF11600
PTF11610
PTF11620
PTF11630
PTF11640
PTF11650
PTF11660
PTF11670
PTF11680
PTF11690
PTF11700
PTF11710
PTF11720
PTF11730
PTF11740
PTF11750
PTF11760

```

```

6000 FORMAT (1H,5X,** ERROR ** RIVER NODE',I4,' IS NOT IN ELEMENT',PTF11770
1I4,' CALCULATED VALUES FOR W1-W3 ARE ',3F10.5) PTF11780
6010 FORMAT (1H,5X,** ERROR ** WELL NODE',I4,' IS NOT IN ELEMENT',IPTF11790
14,' CALCULATED VALUES FOR W1-W3 ARE ',3F10.5) PTF11800
6020 FORMAT (1H,5X,** ERROR ** RIVER NODE',I4,' IS NOT IN ELEMENT',PTF11810
1I4,' CALCULATED VALUES FOR ETA,ZETA ARE ',1P2E19.12) PTF11820
6030 FORMAT (1H,5X,** ERROR ** WELL NODE',I4,' IS NOT IN ELEMENT',IPTF11830
14,' CALCULATED VALUES FOR ETA,ZETA ARE ',1P2E19.12) PTF11840
END PTF11850-

```



```

BND 350
BND 360
BND 370
BND 380
BND 390
BND 400
BND 410
BND 420
BND 430
BND 440
BND 450

```

```

C      LOAD THE TERMS OF THE A MATRIX INTO THE BB MATRIX
C
C      I=0
      DO 160 NB=1,NIJ
      IF (IBNODE(NB).LT.2) GO TO 160
      I=I+1
      BB(1,I)=A(NB,IDIAG)
      IF (NB.EQ.NIJ) GO TO 60
      BB(2,I)=A(NB,IDIAG+1)
      IF (NB+JT-1-NIJ) 20,50,60

```

C

```

BND 460
BND 470
BND 480
BND 490
BND 500
BND 510
BND 520
BND 530
BND 540
BND 550
BND 560
BND 570
BND 580
BND 590
BND 600
BND 610
BND 620
BND 630
BND 640
BND 650
BND 660
BND 670
BND 680
BND 690
BND 700
BND 710

```

```

20 IF (NB+JT-NIJ) 30,40,50
30 BB(3,I)=A(NB,JT+IDIAG+1)
40 BB(4,I)=A(NB,JT+IDIAG)
50 BB(5,I)=A(NB,JT+IDIAG-1)
60 IF (NB.EQ.1) GO TO 160

C      FOR CONFINED CASE, THE TERMS THAT WOULD BE TO THE LEFT OF THE
C      DIAGONAL MUST BE EVALUTED THROUGH THEIR SYMMETRIC COUNTERPART.
C
C      IF (WTABLE) GO TO 110
      BB(6,I)=A(NB-1,2)
      IF (NB-JT) 160,100,70
70 IF (NB-JT-1) 160,90,80
80 BB(7,I)=A(NB-JT-1,JT+2)
90 BB(8,I)=A(NB-JT,JT+1)
100 BB(9,I)=A(NB-JT+1,JT)
      GO TO 160

C      FOR WATER-TABLE CASE, ALL TERMS ARE ON THE SAME LINE OF MATRIX 'A'
C
C      110 BB(6,I)=A(NB,IDIAG-1)
          IF (NB-JT) 160,150,120
120 IF (NB-JT-1) 160,140,130
130 BB(7,I)=A(NB,IDIAG-JT-1)
140 BB(8,I)=A(NB,IDIAG-JT)
150 BB(9,I)=A(NB,IDIAG-JT+1)

```

C

BND1182
 BND1183
 BND1184
 BND1190
 BND1200
 BND1210
 BND1220
 BND1230
 BND1240
 BND1250
 BND1260
 BND1270
 BND1280
 BND1290
 BND1300
 BND1310
 BND1320
 BND1330
 BND1340
 BND1350
 BND1360

```

DO 235 IM=1,12
  TBH=TBH+PBH(IM)*FMON(IM)
  TBF=TBF+PBF(IM)*FMON(IM)
  I=0
DO 340 NB=1,NIJ
  IF (IBNODE(NB).LT.2) GO TO 330
  TQH=TBH*BQH(NB)
C
C   ADJUSTMENT FOR THE ZEROING OUT OF THE COLUMN IN 'A'
C
  IF (WTABLE) GO TO 340
  I=I+1
  IF (NB.EQ.NIJ) GO TO 280
  QRE(NB+1)=QRE(NB+1)-BB(2,I)*TQH
  IF (NB+JT-1-NIJ) 240,270,280
240 IF (NB+JT-NIJ) 250,260,270
250 QRE(NB+JT+1)=QRE(NB+JT+1)-BB(3,I)*TQH
260 QRE(NB+JT)=QRE(NB+JT)-BB(4,I)*TQH
270 QRE(NB+JT-1)=QRE(NB+JT-1)-BB(5,I)*TQH
280 IF (NB.EQ.1) GO TO 340
  QRE(NB-1)=QRE(NB-1)-BB(6,I)*TQH
  
```

C

BND1370
 BND1380
 BND1390
 BND1400
 BND1410
 BND1420
 BND1430
 BND1431
 BND1440
 BND1450
 BND1460
 BND1470
 BND1480
 BND1490
 BND1500
 BND1510

```

  IF (NB-JT) 340,320,290
290 IF (NB-JT-1) 320,310,300
300 QRE(NB-JT-1)=QRE(NB-JT-1)-BB(7,I)*TQH
310 QRE(NB-JT)=QRE(NB-JT)-BB(8,I)*TQH
320 QRE(NB-JT+1)=QRE(NB-JT+1)-BB(9,I)*TQH
  GO TO 340
330 IF (IBNODE(NB).LT.1) GO TO 340
  TQH=TBF*BQH(NB)
  QRE(NB)=QRE(NB)+TQH
  IF (TQH.GT.0.0) BUD(3) = BUD(3) + TQH
  IF (TQH.LT.0.0) BUD(4) = BUD(4) + TQH
340 CONTINUE
DO 350 NB=1,NIJ
  IF (IBNODE(NB).LT.2) GO TO 350
  QRE(NB)=0.
350 CONTINUE
  RETURN
  END
  
```



```

QBQ=BB(1,I)*H(NB)
IF (NB.EQ.NIJ) GO TO 420
QBQ=QBQ+BB(2,I)*H(NB+1)
IF (NB+JT-1-NIJ) 380,410,420
380 IF (NB+JT-NIJ) 390,400,410
390 QBQ=QBQ+BB(3,I)*H(NB+JT+1)
400 QBQ=QBQ+BB(4,I)*H(NB+JT)
410 QBQ=QBQ+BB(5,I)*H(NB+JT-1)
420 IF (NB.EQ.1) GO TO 470
QBQ=QBQ+BB(6,I)*H(NB-1)
IF (NB-JT) 470,460,430
430 IF (NB-JT-1) 470,450,440
440 QBQ=QBQ+BB(7,I)*H(NB-JT-1)
450 QBQ=QBQ+BB(8,I)*H(NB-JT)
460 QBQ=QBQ+BB(9,I)*H(NB-JT+1)
470 QB(NB)=QBQ-QB(NB)
IF(NN.EQ.NSTEP) GO TO 472
IF(MOD(NN,NPRNT).EQ.0) GO TO 472
IF(NN.EQ.1) GO TO 472
MPRNT = NPRNT/2
IF(MOD(NN,NPRNT).EQ.0) GO TO 472
GO TO 475
472 WRITE(IO,6050) NB, QB(NB)
6050 FORMAT(1H,1X,'NODE=',I3,1X,'SPEC HEAD IN-OUT',F10.5)
475 IF (QB(NB).GT.0.0) QSH = QSH + QB(NB)
IF (QB(NB).LT.0.0) QSHM = QSHM + QB(NB)
480 CONTINUE
IF(.NOT. LRIVER)GO TO 488
IF(NN.EQ.NSTEP)GO TO 485
IF(MOD(NN,NPRNT).EQ.0) GO TO 485
IF(NN.EQ.1)GO TO 485
IF(MOD(NN,NPRNT).EQ.0)GO TO 485
GO TO 488
485 WRITE(IO,6068)
6068 FORMAT(1H0,5X,'RIVER-NODE FLUXES'//)
DO 486 I=1,32
II=I+32
486 WRITE(IO,6070) I,QRR(I),II,QRR(II)

```

```

BND1710
BND1720
BND1730
BND1740
BND1750
BND1760
BND1770
BND1780
BND1790
BND1800
BND1810
BND1820
BND1830
BND1840
BND1850
BND1860
JEC
JEC
JEC
JEC
JEC
JEC
JEC
JEC
BND1870
BND1875
BND1880
JEC
JEC
JEC
JEC
JEC
JEC
JEC
JEC
JEC
JEC

```

JEC
JEC
JEC

BND1890
BND1895
BND1900
BND1910
BND1920
BND1930
BND1940
BND1950
BND1960
BND1970

JEC
BND1980
BND1990
BND2000

```

6070  FORMAT(1X,I3,F12.5,1X,I3,F12.5)
      WRITE(IO,6073)
6073  FORMAT(1H,/)
488   CONTINUE
C
      BUD(1)=QSH
      BUD(2)=QSHM
      IF (STEADY) GO TO 500
      BUD5=0.0
      DO 490 I=1,NIJ
490    BUD5=BUD5+3.0*UV(I)*H(I)
      BUD(5)=BUD(5)-BUD5
500    CONTINUE
      BALMAS=0.
      BALTOT=0.
      WRITE(IO,7000)
7000  FORMAT(1H,/,6X,'BUDGET BALANCE'/)
      WRITE(IO,8000)
8000  FORMAT(1H,28X,'FT**3/S ',5X,'FT**3',/)
      DO 510 L=1,NBUD
510    BUDTOT(L)=BUDTOT(L)+BUD(L)*DELT
      BALMAS=8ALMAS+BUD(L)

```

C

BND2010
BND2020
BND2030
BND2040
BND2050
OUT 220
OUT 230
OUT 240
OUT 250

```

      BALTOT=BALTOT+BUDTOT(L)
      WRITE (IO,6010) (BUDTTL(N,L),N=1,5),BUD(L),BUDTOT(L)
510  BUD(L)=0.
      WRITE (IO,6020)
      WRITE (IO,6030) BALMAS,BALTOT
      DO 600 I=1,IS
      K=(I-1)*JT
      DO 610 J=1,JT
      HI(K+J)=OLDH(K+J)-H(K+J)
      SATTH = T(K+J)/.0015
      IF (SATTH .GE. HI(K+J)) GO TO 610
      WRITE (IO,6005) I,J
      C 610 CONTINUE
      C 600 CONTINUE
      C 6005 FORMAT(1H,10X,'NODE',I3,',',I3,' GOES DRY')
      IF (.NOT.WTABLE) RETURN

```

BND2060


```

DO 10 K=1,IACCUM
I=N-K
J=L+K
10 SUM=SUM-A(I,K+1)*A(I,J)
20 IF (L.GT.1) GO TO 40
IF (SUM.LE.0) GO TO 30
ADIAGN=1./DSQRT(SUM)
A(N,L)=ADIAGN
GO TO 50
30 WRITE (IO,6000) N
STOP2
40 A(N,L)=SUM*ADIAGN
50 CONTINUE
RETURN
6000 FORMAT (1H1,5X,24HDECOMPOSITION FAILS ROW ,I3,16H HAS ZERO LENGTH)SDC 860
END-----SDC 260
SDC 270
SDC 280
SDC 290
SDC 300
SDC 310
SDC 320
SDC 330
SDC 340
SDC 350
SDC 360
SDC 370
SDC 380
SDC 390
SDC 400
SDC 410
SUBROUTINE SSOLVE(A,NIJ,NBW,B,X)-----SDC 430
-----SDC 440
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCSDC 450
C THIS ENTRY DOES THE FORWARD AND BACKWARD CC
C SUBSTITUTIONS NEEDED TO SOLVE THE SYSTEM OF CC
C EQUATIONS AND RETURNS THE NEW HEADS IN 'X' CC
C CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCSDC 460
C IMPLICIT DOUBLE PRECISION (A-H,O-Z) SD
DIMENSION A(NIJ,NBW),B(1),X(1) SD
COMMON /IOUNIT/ IN,IO,IPRINT,NPRNT SD
FORWARD SUBSTITUTE FOR LOWER TRIANGULAR SOLUTION SD
C SD
C SD
C SD

```

```

C -----
C B - THE R.H.S. OF THE EQUATION
C X - THE UNKNOWN VECTOR
C -----
C
C      DO 80 N=1,NIJ
C      SUM=B(N)
C      IB=N-NBW+1
C      IF (IB.LT.1) IB=1
C      IE=N-1
C      IF (IB.GT.IE) GO TO 70
C      DO 60 K=IB,IE
C      I=N-K+1
C      60 SUM=SUM-A(K,I)*X(K)
C      70 X(N)=SUM*A(N,1)
C      80 CONTINUE
C
C      BACKWARD SUBSTITUTE FOR UPPER TRIANGULAR SOLUTION
C
C      DO 110 N=1,NIJ
C      IJ=NIJ-N+1
C      IE=IJ+NBW-1
C      IF (IE.GT.NIJ) IE=NIJ
C      SUM=X(IJ)
C
C      IB=IJ+1
C      IF (IB.GT.IE) GO TO 100
C      DO 90 K=IB,IE
C      I=K-IJ+1
C      90 SUM=SUM-A(IJ,I)*X(K)
C      100 X(IJ)=SUM*A(IJ,1)
C      110 CONTINUE
C      RETURN
C      END
C -----
C
C      SDC 550
C      SDC 560
C
C      SDC 570
C      SDC 580
C      SDC 590
C      SDC 600
C      SDC 610
C      SDC 620
C      SDC 630
C      SDC 640
C      SDC 650
C      SDC 660
C      SDC 670
C      SDC 680
C      SDC 690
C      SDC 700
C      SDC 710
C      SDC 720
C      SDC 730
C      SDC 740
C      SDC 750
C      SDC 760
C
C      SDC 770
C      SDC 780
C      SDC 790
C      SDC 800
C      SDC 810
C      SDC 820
C      SDC 830
C      SDC 840
C      SDC 850
C      SDC 870-

```



```

70 XX(I)=B(I)+SUM
   XX(NIJ)=XX(NIJ)/A(NIJ,IHBP)
   DO 90 IBACK=2,NIJ
   I=NN-IBACK
   JP=I
   KR=IHBP+1
   MR=MINO(NBW,IHALFB+IBACK)
   SUM=0.0
   DO 80 J=KR,MR
   JP=JP+1
80 SUM=SUM+A(I,J)*XX(JP)
90 XX(I)=(XX(I)-SUM)/A(I,IHBP)
   RETURN
   END
-----
C SUBROUTINE QZEE (UV)
C -----
C IMPLICIT DOUBLE PRECISION (A-H,O-Z)
C REAL BUDTTL
C LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLLOT,EVPTRN
C COMMON/AQUIFR/ H(572),HP(572),QRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
C COMMON/WTBLE/ QBAR(572),BOTT(572),HO(572)
C COMMON /BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
C COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYP
C COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KDYS,IDYR,MON(12)MAI
C COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLLOT,EVPTRN
C COMMON/WELLS/ QW(200),AREAPP(200),QAPP(1020),REAPP(1020),
1WW(4,200),PMPDMD(5,12),PMPAPP(5),NEW(200),NTW(200),NAPP(200),
2NEAPP(30,200),NWELLS,NSEWER,NWTYP
C COMMON /WETHER/ PRECIP(12),U(12),SC,DEPTH,SF,ETFAC(8)
C COMMON /MPSCCL/ XSCL,YSCCL,XO,YO
C COMMON /IOUNIT/ IN,IO,IPRINT,NPRNT

```

```

ADC 670
ADC 680
ADC 690
ADC 700
ADC 710
ADC 720
ADC 730
ADC 740
ADC 750
ADC 760
ADC 770
ADC 780
ADC 790
ADC 800-
QZE 10
QZE 20
QZE 30
QZE 50
QZE 60
QZE 70
QZE 80
QZE 90
QZE 140
QZE 110
QZE 120
QZE 130
QZE 140
QZE 150
QZE 160
QZE 170

```

C

```

DIMENSION UV(1),LAT(3),LON(3)
QZE 180
QZE 190

-----
PRECIP - MONTHLY VALUES FOR PRECIPITATION
QZE 200
U - MONTHLY VALUES FOR E.T. DEMAND
QZE 210
SC - SOIL MOISTURE STORAGE TERM
QZE 220
DEPTH - EXTINCTION DEPTH FOR E.T.
QZE 230
SF1-SF4 - COEFFICIENTS USED TO CALCULATE SOIL MOISTURE E.T. FACTOR
QZE 240
ETFAC - SOIL MOISTURE E.T. FACTORS
QZE 250

-----

READ (IN,5010) (PRECIP(M),M=1,12)
QZE 260
IF (IPRINT.GE.2) WRITE (IO,6010) (PRECIP(M),M=1,12)
QZE 270
IF (.NOT.EVPTRN) RETURN
QZE 280
READ (IN,5010) (U(M),M=1,12)
QZE 290
IF (IPRINT.GE.2) WRITE (IO,6020) (U(M),M=1,12)
QZE 300
READ (IN,5010) SC,DEPTH
QZE 310
IF (IPRINT.GE.2) WRITE (IO,6030) SC,DEPTH
QZE 320
SF=ETFAC(1)+SC*(ETFAC(2)+SC*(ETFAC(3)*SC-ETFAC(4)))
QZE 330
RETURN
QZE 340
QZE 350
5010 FORMAT (12F6.0)
QZE2240
6010 FORMAT (1H1/1H ,30X,'PRECIPITATION AND EVAPOTRANSPIRATION DATA',1HQZE2280
10,10X,'MONTHLY PRECIPITATION VALUES',1H ,5X,12F8.3)
QZE2290
6020 FORMAT (1H0,10X,'MONTHLY VALUES FOR E.T. DEMAND',1H ,12F10.4)
QZE2300
6030 FORMAT (1H0,5X,'SOIL MOISTURE STORAGE TERM = ',F10.4/1H ,5X,'EXTINQZE2310
CTION DEPTH FOR E.T. = ',F10.4)
QZE2320
END
QZE 360
C

```

```

-----
SUBROUTINE PUMPNU (UV)
-----
      IMPLICIT DOUBLE PRECISION (A-H,O-Z)
      REAL BUDTTL
      LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLLOT,EVPTRN
      COMMON/AQUIFR/ H(572),HP(572),QRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
      COMMON/WTBLE/ QBAR(572),BOTT(572),HO(572)
      COMMON /BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
      COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYP
      COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KDYS,IDYR,MON(12)
      COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLLOT,EVPTRN
      COMMON/WELLS/ QW(200),AREAPP(200),QAPP(1020),REAPP(1020),
1WW(4,200),PMPDMD(5,12),PMPAPP(5),NEW(200),NTW(200),NAPP(200),
2NEAPP(30,200),NWELLS,NSEWER,NWTYP
      COMMON /WETHER/ PRECIP(12),U(12),SC,DEPTH,SF,ETFAC(8)
      COMMON /MPSCCL/ XSCL,YSCCL,XO,YO
      COMMON /IOUNIT/ IN,IO,IPRINT,NPRNT
      DIMENSION UV(1),LAT(3),LON(3)
-----
      NWELLS - TOTAL NUMBER OF WELLS
      NSEWER - NO. OF WELLS WHOSE OUTFLOW CONTRIBUTES DIRECTLY TO THE
      NWTYP - NO. OF WELL TYPES
      PMPAPP - PERCENT OF PUMPAGE THAT IS APPLIED WATER BY PUMPAGE TYPE
      PMPDMD - MONTHLY PUMPAGE DEMAND DISTRIBUTION BY PUMPAGE TYPES
-----
      READ (IN,5000) NWELLS,NSEWER,NWTYP
      IF (IPRINT.GE.2) WRITE (IO,6040) NWELLS,NSEWER,NWTYP

      IF (NWELLS.LT.1) RETURN

      IF (IPRINT.GE.2) WRITE (IO,6050)
      DO 20 N=1,NWTYP
      READ (IN,5011) (PMPDMD(N,M),M=1,12),PMPAPP(N)
      IF (IPRINT.GE.2) WRITE (IO,6060) N,PMPAPP(N)
      IF (IPRINT.GE.2) WRITE (IO,6070) (PMPDMD(N,M),M=1,12)
      IF (IPRINT.GE.4) WRITE (IO,6080)
20

```



```

C      CALCULATE LOCAL WELL COORDINATES WITHIN THE ELEMENT
C
      CALL PTFNDR (XW,YW,NEW(N),WW,N,2)
      70 CONTINUE
      WRITE(10,7000)
      7000 FORMAT(1H ,120('---') )
C
      RETURN
      5000 FORMAT (20I4)
      5011 FORMAT (13F5.0)
      5020 FORMAT (9I4,4X,F10.0)
      5030 FORMAT (2F12.0,3I4,4X,F10.0)

C
      6000 FORMAT (1H ,5X,I4,2F12.3,2X,I5,2X,I5,2X,I5,2X,F12.4)
      6040 FORMAT (1H1/1H ,30X,'WELL DATA'/1H0,5X,'NUMBER OF WELLS (NWELLS)',QZE2270
      1' = ',I4/1H ,5X,'NUMBER OF SEWERED WELLS (NSEWER) = ',I4/1H ,5X,'NQZE2330
      2NUMBER OF WELL TYPES (NWTYPE) = ',I2)
      6050 FORMAT (1H0,30X,'WELL TYPE PARAMETERS')
      6060 FORMAT (1H0,10X,'TYPE ',I4/1H ,5X,'% OF PUMPAGE APPLIED =',F6.3)
      6070 FORMAT (1H ,10X,'MONTHLY PUMPAGE DEMAND VALUES'/1H ,5X,12F10.3)
      6080 FORMAT (1H0,30X,'WELL DATA'/1H0,5X,'WELL',3X,'X-COORD.',5X,'Y-COORQZE2380
      1D.',3X,'ELEM.',2X,'TYPE',2X,'# APPLIED',3X,'WITHDRAWL'/)
      6090 FORMAT (1H ,30X,'WATER IS APPLIED ON ELEMENTS',2X,10I5,
      $ (11X,20I5))
      END
C
      QZE 930
      QZE 940
      QZE 950
      QZE 960

      QZE 970
      QZE 980
      QZE2230
      QZE2241
      QZE2250
      QZE2260

      QZE 990

```

```

C -----
C SUBROUTINE LOADQR (UV)
C -----
C
C IMPLICIT DOUBLE PRECISION (A-H,O-Z)
C REAL BUDTTL
C LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
C COMMON/AQUIFR/ H(572),HP(572),GRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
C COMMON/WTBLE/ QBAR(572),BOTT(572),HO(572)
C COMMON /BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
C COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYP
C COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KDYS,IDYR,MON(12)MAI 140
C COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN QZE 110
C COMMON/WELLS/ QW(200),AREAPP(200),QAPP(1020),REAPP(1020),
C 1W(4,200),PMPDMD(5,12),PMPAPP(5),NEW(200),NTW(200),NAPP(200),
C 2NEAPP(30,200),NWELLS,NSEWER,NWTYP QZE 130
C COMMON /WETHER/ PRECIP(12),U(12),SC,DEPTH,SF,ETFAC(8) QZE 140
C COMMON /MPSCCL/ XSCL,YSCCL,XO,YO QZE 150
C COMMON /IOUNIT/ IN,IO,IPRINT,NPRNT QZE 160
C DIMENSION UV(1),LAT(3),LON(3) QZE 170
C ZERO THE ELEMENT APPLICATION AND RECHARGE ARRAYS QZE 180
C
C DO 80 NE=1,NEL QZE1010
C QAPP(NE)=0. QZE 20
C 80 REAPP(NE)=0.
C
C LOAD THE SEWERED PUMPAGE (RETURN FLOW DIRECTLY TO RIVER)
C
C ASEWR=0.
C QSEWR=0.
C EXCESS=0.
C IF (NWELLS.LT.1) GO TO 170
C IF (NSEWER.LT.1) GO TO 130
C DO 100 NS=1,NSEWER
C ASEWR=ASEWR+AREAPP(NS)
C Q=PMPDMD(NTW(NS),KMN)*QW(NS)
C IF (Q.GT.0.0) BUD(7) = BUD(7) - Q
C IF (Q.LT.0.0) BUD(6) = BUD(6) - Q
C QSEWR=QSEWR+Q
C
C QZE1010
C QZE 20
C
C QZE 30
C QZE 50
C QZE 60
C QZE 70
C QZE 80
C QZE 90
C QZE 140
C QZE 110
C QZE 120
C QZE 130
C QZE 140
C QZE 150
C QZE 160
C QZE 170
C QZE 180
C QZE1020
C QZE1030
C QZE1040
C QZE1050
C QZE1060
C QZE1070
C QZE1080
C QZE1090
C QZE1100
C QZE1110
C QZE1120
C QZE1130
C QZE1140
C QZE1150
C QZE1160
C QZE1170
C QZE1180
C QZE1185
C QZE1190

```

QZE1200
QZE1210
QZE1220
QZE1230
QZE1240
QZE1250
QZE1260

```

EXCESS=EXCESS+Q-PMPAPP(1)*QW(NS)
NER=NEW(NS)
DO 90 I=1,NETYP
  NI=NGL(I,NER)
  90 QRE(NI)=QRE(NI)-Q*WW(I,NS)
  100 CONTINUE
  SEWAGE=QSEWR

```

C

QZE1270
QZE1280
QZE1290
QZE1300

```

IF (EXCESS.LT.0.) GO TO 130
SEWAGE=SEWAGE-EXCESS
C LOAD THE APPLIED SPRINKLER WATER INTO THE APPLIED ELEMENTS
BUD(8)=BUD(8)+EXCESS
IF(ASEWR .EQ. 0.) GO TO 106
QA=EXCESS/ASEWR

```

QZE1310

106

```

QA = 0.0
DO 120 NS=1,NSEWER
  NAPPY=NAPP(NS)
  IF (NAPPY.EQ.0) GO TO 120
  DO 110 L=1,NAPPY
    NL=NEAPP(L,NS)
    110 QAPP(NL)=QAPP(NL)+QA
  120 CONTINUE
  130 CONTINUE

```

QZE1320
QZE1330
QZE1340
QZE1350
QZE1360
QZE1370
QZE1380
QZE1390
QZE1400
QZE1410
QZE1420
QZE1430
QZE1440
QZE1450
QZE1455
QZE1460
QZE1470
QZE1480
QZE1490
QZE1500*
QZE1510

```

C LOAD OTHER PUMPAGE TYPES
NSEWR1=NSEWR+1
IF (NSEWR1.GE.NWELLS) GO TO 170
DO 160 NS=NSEWR1,NWELLS
  Q=PMPDMD(NTW(NS),KMN)*QW(NS)
  IF (Q.GT.0.0) BUD(7) = BUD(7) - Q
  IF (Q.LT.0.0) BUD(6) = BUD(6) - Q
  NER=NEW(NS)
DO 140 I=1,NETYP
  NI=NGL(I,NER)
  140 QRE(NI)=QRE(NI)-Q*WW(I,NS)
C IF (NAPP(NS).LT.1) GO TO 160

```



```

C      LOAD THE APPLIED AND RECHARGED WATER INTO THE ELEMENTS
QA=Q*PMPAPP(NTW(NS))
QR=Q-QA
BUD(8)=BUD(8)+QA
BUD(9)=BUD(9)+QR
QA=QA/AREAPP(NS)
QR=QR/AREAPP(NS)
NAPPLY=NAPP(NS)
DO 150 L=1,NAPPLY
NL=NEAPP(L,NS)
REAPP(NL)=REAPP(NL)+QR
150 QAPP(NL)=QAPP(NL)+QA
160 CONTINUE

C      APPLY RAIN AND CALCULATE ET FOR EACH ELEMENT
C      -----
C      RAIN - MONTHLY PRECIP VALUE IN FT/SEC
C      UET - MONTHLY VALUE FOR E.T. DEMAND
C      -----
C      CALCULATED E.T.
C
C      170 RDT=1./(IDMON(KMN)*86400.)
RAIN=PRECIP(KMN)*RDT/12.
IF (EVPTRN) UET=U(KMN)*RDT/12.
IF (EVPTRN) ETFAC=SF*(10.**((ETFAC(5)*U(KMN)))*RDT/12.
DO 220 NE=1,NEL
APPLY=QAPP(NE)+RAIN
BUD(10)=BUD(10)+RAIN*AREA(NE)
IF (.NOT.EVPTRN) GO TO 200
ET=ETFAC*(ETFAC(6))*(12.*APPLY/RDT)**ETFAC(7)-ETFAC(8))
IF(ET.LT.0.0) ET=0.0
C
C      IF (ET.GT.UET) ET=UET
C      IF (ET.GT.APPLY) ET=APPLY
C      APPLY=APPLY-ET
C      BUD(11)=BUD(11)-ET*AREA(NE)
C      IF (ET.GE.UET) GO TO 200

```

```

C      COMPUTE SURFACE INDUCED ET FOR EACH ELEMENT
      DET=UET-ET
      AR=-DEPTH
      DEPSAT=0.0
      DO 180 I=1,NETYP
        NI=NGL(I,NE)
        180 DEPSAT=DEPSAT+DMAX1((H(NI)-ALSD(NI)),AR)
        DEPSAT=DEPSAT/(1.*NETYP)
        IF (DEPTH+DEPSAT) 200,200,190
        190 ADET=DET*(DEPTH+DEPSAT)/DEPTH
        IF (ADET.GT.DET) ADET=DET
        APPLY=APPLY-ADET
        BUD(12)=BUD(12)-ADET*AREA(NE)
        LOAD THE ELEMENT RECHARGE INTO THE NODAL RECHARGE ARRAY
        200 QR=(APPLY+REAPP(NE))*AREA(NE)/(1.*NETYP)
        DO 210 I=1,NETYP
          NI=NGL(I,NE)
          210 QRE(NI)=QRE(NI)+QR
          220 CONTINUE
          IF (STEADY) RETURN
          BUD5=0.0
          DO 230 I=1,NIJ
            C
            C      CALCULATE TIME DERIVATIVE BASED ON PAST HEAD VALUES
            C      AND UPDATE THE PAST HEADS
            C
            STOR=UV(I)*(4.*H(I))-HP(I)
            C
            C      CORRECT TIME DERIVATIVE TERM FOR WATER-TABLE
            C
            IF (WTABLE) STOR=STOR-UV(I)*3.0*H0(I)
            HP(I)=H(I)
            BUD5=BUD5+STOR
            230 QRE(I)=QRE(I)+STOR
            BUD(5)=BUD5
            RETURN
            END
      END

```

QZE1860
 QZE1870
 QZE1880
 QZE1890
 QZE1900
 QZE1910
 QZE1920
 QZE1930
 QZE1940
 QZE1950
 QZE1960
 QZE1970
 QZE1980
 QZE1990
 QZE2000
 QZE2010
 QZE2020
 QZE2030
 QZE2040
 QZE2050
 QZE2060
 QZE2070
 QZE2080
 QZE2090
 QZE2100
 QZE2110
 QZE2120
 QZE2130
 QZE2140
 QZE2150
 QZE2160
 QZE2170
 QZE2180
 QZE2190
 QZE2200
 QZE2210
 QZE2420-

```

C -----
SUBROUTINE OUTPUT (IPR)
C -----
IMPLICIT DOUBLE PRECISION (A-H,O-Z)
REAL XSCLP,YSCLP
LOGICAL STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON/AQUIFR/ H(572),HP(572),QRE(572),QZ(572),T(572),S(572),
1ALSD(572),X(572),Y(572),AREA(1020),SAREA
COMMON/RIVERB/ HPOOL(2),STAGE(20,2),DSCHRG(20,2),QCOMR(118),
1ELVR(118),XFACTR(118),RW(4,118),RIVERH(118),
2RIVERQ(118),FLOWIN(366,7),QCOR(5),SEWAGE,NODEIN(7),NSTAGE(2),
3NRN(118),NER(118),NRTYP(118),NREACH,NRNODE,NTRIB,QRR(64)
COMMON /SIZE/ IS,JT,NIJ,NEL,NBW,IDIAG,NGL(4,1020),NETYP
COMMON /LOGIC/STEADY,WTABLE,QUAD,LATLON,LRIVER,DRAWD,LPLOT,EVPTRN
COMMON /IUNIT/ IN,IO,IPRINT,NPRNT
COMMON /PLT/ XSCLP,YSCLP,IPLOT
COMMON/HEAD2/ HI(572),OLDH(572)
COMMON/TIME2/ DELT,BETA,EPS,NSTEP,MAXIT,NN

WRITE (IO,6020)
WRITE(IO,6035)
DO 10 I=1,IS
K=(I-1)*JT
IKK = K + 1
WRITE(IO,6000) I,IKK,(H(K+J),J=1,JT)
10 CONTINUE
IF (.NOT.DRAWD) GO TO 40
WRITE (IO,6030)
WRITE(IO,6035)
DO 30 I=1,IS
K=(I-1)*JT
DO 20 J=1,JT
HI(K+J)=OLDH(K+J)-H(K+J)
IKK = K + 1
20 WRITE(IO,6000) I,IKK,(HI(K+J),J=1,JT)
25 CONTINUE
30 CONTINUE
40 CONTINUE

```

```

IF (LRIVER) WRITE (IO,6040)
IF (LRIVER) WRITE (IO,6010) (N,RIVERH(N),ELVR(N),RIVERQ(N),N=1,
1NRNODE)
C IF (LPLT.AND.IPLT2.EQ.9) CALL GPCP(H,NIJ,' HEAD ','VALUES')
C IF (LPLT.AND.IPLT2.EQ.10) CALL VPLOT
RETURN
6000 FORMAT(1H,2X,I4,2X,I4,2X,(10F10.2))
6010 FORMAT (5X,I4,3F10.2)
6020 FORMAT (1H0,30X,'HEAD VALUES'/)
6030 FORMAT (1H0,30X,'DRAWDOWNS FROM INITIAL HEAD VALUES'/)
6035 FORMAT(1H,2X,' ROW',2X,'NODE')
6040 FORMAT (1H0,30X,'RIVER-STAGE AND DISCHARGE VALUES'/,1H,5X,'NODE',OUT 390
13X,'STAGE',4X,'BOTTOM ELV',4X,'DISCHARGE'/)
6111 FORMAT(5F8.2)
END
C -----
C BLOCK DATA
C -----
IMPLICIT DOUBLE PRECISION (A-H,O-Z)
REAL BUDTTL,TITLE1,TITLE2,XSCLP,YSCLP
COMMON/BNDRY/BQH(572),QB(572),PBH(12),PBF(12),BB(9,50),IBNODE(572)BLK 22
COMMON /BUDGET/ BUDTTL(5,20),BUD(20),BUDTOT(20),NBUD
COMMON /IUNIT/ IN,IO,IPRINT,NPRNT
COMMON /MPACL/ XSCL,YACL,XO,YO
COMMON /PLT/ XSCLP,YSCLP,IPLT
COMMON /TIME/FMON(12),IDMON(12),KYR,KMN,KDY,IDAY,KDYS,IDYR,MON(12)BLK 70
COMMON/TIME2/ DELT,BETA,EPS,NSTEP,MAXIT
COMMON /WETHER/ PRECIP(12),U(12),SC,DEPTH,SF,ETFAC(8)
DIMENSION TITLE1(5,12),TITLE2(5,6)
EQUIVALENCE (BUDTTL(1,1),TITLE1(1,1)),(BUDTOT(1,1),TITLE2(1,1))
DATA IN/15/,IO/16/
DATA XSCLP/1.0/,YSCLP/1.0/,IPLT/0/
DATA XSCL/1.0/,YACL/1.0/,XO/1.E30/,YO/1.E30/
DATA U/12*0.0/,SC/0.0/,DEPTH/0.0/
DATA BETA/1.0/,EPS/0.0/,MAXIT/1/
DATA NBUD/12/
DATA ETFAC/0.531747,0.295164,0.003804,0.057697,
1 0.02426,0.70917,0.82416,0.11556/
DATA BB/450*0.0/
DATA TITLE1/240H SPEC. HEAD INFLOW SPEC. HEAD OUTFLOW SPEC. BDBLK 180
1RY INFLOW SPEC. BDRY OUTFLOW STORAGE CHANGE PUMP INJECTN WBLK 190
2ELLS PUMP DISCHRG WELLS APPLIED WATER DIRECT RECHARGE
BLK 200

```

3PRECIPITATION	EVAPOTRNSPIRATION	GROUND-WATER E.T.	BLK	205
DATA TITILE2/120H	1ST REACH LOSS	2ND REACH LOSS	3RD REACH LOSS	210
1CH LOSS	4TH REACH LOSS	5TH REACH LOSS	6TH REACH LOSS	220
2S /				BLK 230
DATA BUD/20*0./,BUDTOT/20*0./				BLK 270
DATA FMON/12*0.0/				
DATA IDMON/ 31,28,31,30,31,30,31,30,31,30,31/				BLK 280
DATA MON/ "JAN","FEB","MAR","APR","MAY","JUN",				BLK 290
1 "JUL","AUG","SEP","OCT","NOV","DEC" /				BLK 300
END				BLK 310-

Model Input for Example Problem

```

00001:EXAMPLE SETUP PROBLEM: ONE RIVER,ONE INJECTION WELL,ONE PUMPING WELL,
00002:      ET AND PRECIPITATION OPTION, AND CONFINED AQUIFER
00003:      F      F      F      F      T      T      T      T
00004:      6      1
00005:1983 01 01 1 5
00006: 20      1.0      .0001
00007:      1.00      1.00
00008: 6      4000.0      4000.0
00009: 20 20
00010: 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2
00011: 0
00012: 0
00013: 0
00014: 0
00015: 0
00016: 0
00017: 0
00018: 0
00019: 0
00020: 0
00021: 0
00022: 0
00023: 0
00024: 0
00025: 0
00026: 0
00027: 0
00028: 0
00029: 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1
00030: HEAD 0 ( 9F6.0) 200.0 1
00031: HP 0 ( 9F6.0) 200.0 1
00032: STOR 0 ( 9F6.0) 1.000E-2 1
00033: TRAN 0 ( 9F6.0) 1.15740E-1 1
00034: ALSD 0 ( 9F6.0) 200.0 1
00035: BQM 15 (20F4.0) 1.00 1
00036: 200 200 200 200 200 200 200 200 200 200 200 200 200 200 200 200 200 200
00037:
00038:
00039:
00040:
00041:
00042:
00043:
00044:
00045:
00046:
00047:
00048:
00049:
00050:
00051:
00052:
00053:
00054:
00055: .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01 .01
00056: 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0
00057: 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0
00058: 0.0 0.0 0.0 0.0 19000.0
00059: 1000.0 0.0 1000.0 19000.0

```


Model Output for Example Problem

(partial listing)

1

1

```

-----
C          EXAMPLE SETUP PROBLEM: ONE PIVER, ONE INJECTION WELL, ONE PUMPING WELL,
          ET AND PRECIPITATION OPTION, AND CONFINED AQUIFER
C-----
          MODEL RUN DATE:   WED, OCT 26 1983
          MODEL RUN TIME:   13:32:37
C-----

```

```

C          LOGICAL CONTROL VARIABLES
          STEADY   WTABLE   QUAD   LATLON   LRIVER   DRAWD   LPLOT   EVPTRN
          F         F         F         F         T         T         T         T
C          THE PRINTOUT CONTROL PARAMETER (IPRINT) =      6
          NUMBER OF TIME STEPS BETWEEN EXTENDED PRINTOUTS (NPRNT) =      1
C          THE STARTING DATE FOR THE SIMULATION IS : JAN  1, 1983
          NUMBER OF DAYS PER TIME STEP (IDAY) =      1
          NUMBER OF TIME STEPS TO BE TAKEN (NSTEP) =      5
C          ITERATION PARAMETERS FOR STREAMFLOW SOLUTION ROUTINE
          MAXIMUM NUMBER OF ITERATIONS (MAXIT) =      20
          CONVERGENCE ACCELERATION PARAMETER (BETA) =  1.00000
          CONVERGENCE CRITERION (EPS) = 0.000100
C          SCALING FACTOR FOR X - COORDINATES (XSCL) =      1.000
          SCALING FACTOR FOR Y - COORDINATES (YSCL) =      1.000
C          PLOT CONTROL PARAMETER (IPLOT) =      6
          PLOT SCALING FACTOR IN X DIRECTION (XSCLP) =  4000.000
          PLOT SCALING FACTOR IN Y DIRECTION (YSCLP) =  4000.000
C          NUMBER OF NODES IN ROW =      20
          NUMBER OF NODES IN COLUMN =      20
C          NUMBER OF NODES (NIJ) =      400
          NUMBER OF ELEMENTS (NEL) =      722
          HALF BANDWIDTH (NBW) =      22

```

```

1          BOUNDARY-CONDITION TYPE MAP
          0 - INTERIOR NODE
          1 - SPECIFIED-FLUX BOUNDARY NODE
          2 - SPECIFIED-HEAD BOUNDARY NODE

```

ROW	NODE	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15	16	17	18	19	20
1	1	2	2	2	2	2	2	2	2	2	2	2	2	2	2	2	2	2	2	2	2
2	21	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
3	41	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
18	341	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
19	361	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0
20	381	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1

```

1          INPUT DATA FOR ARRAY   HEAD
0          HEAD WAS NOT READ, ENTIRE MATRIX SET EQUAL TO 2.000E+02
1          INPUT DATA FOR ARRAY   HP
0          HP WAS NOT READ, ENTIRE MATRIX SET EQUAL TO 2.000E+02
1          INPUT DATA FOR ARRAY   STOR
0          STOR WAS NOT READ, ENTIRE MATRIX SET EQUAL TO 1.000E-02
1          INPUT DATA FOR ARRAY   TRAN
0          TRAN WAS NOT READ, ENTIRE MATRIX SET EQUAL TO 1.157E-01
1          INPUT DATA FOR ARRAY   ALSO
0          ALSO WAS NOT READ, ENTIRE MATRIX SET EQUAL TO 2.000E+02
1          INPUT DATA FOR ARRAY   BQH

```

```

1
0      EQM WAS READ FROM UNIT 15, CN FORMAT      (20F4.0) , AND SCALES BY      1.000E+00

      ROW NODE
      1 1      2.000E+02 2.000E+02 2.000E+02 2.000E+02 2.000E+02 2.000E+02 2.000E+02 2.000E+02 2.000E+02
      2 2 21 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01
      0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01
      19 361 C.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01
      C.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01 0.000E-01
      20 381 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02
      1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02 1.000E-02
      0      1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00
      0      1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00 1.00
      1      THE X AND Y OFFSETS FOR THE GRID ORIGIN ARE      0.0000      0.0000
      0      NODAL X - Y COORDINATES
      1 0.0000
      2 0.0000 1000.0000
      3 0.0000 2000.0000
      4 0.0000 3000.0000
      397 19000.0000 16000.0000
      398 19000.0000 17000.0000
      399 19000.0000 18000.0000
      400 19000.0000 19000.0000
      1      GLOBAL INDICES AND ELEMENT AREAS
      1 1 21 2 500000.0000
      2 2 21 22 500000.0000
      3 2 22 3 500000.0000
      4 3 22 23 500000.0000
      719 378 398 379 500000.0000
      720 379 398 399 500000.0000
      721 379 399 380 500000.0000
      722 380 399 400 500000.0000
      TOTAL MODEL AREA = 3.61000E+08
      1      RIVER PARAMETERS
      2      NUMBER OF RIVER REACHES = 1
      3      NUMBER OF RIVER NODES = 18
      4      NUMBER OF RIVER-LEAKAGE ZONES = 1
      0      RIVER-LEAKAGE VALUES BY ZONE (QCOR(N))
      1      REACH NUMBER 1
      2      RIVER-REACH PARAMETERS
      3      3 POINTS USED IN STAGE - DISCHARGE RELATIONSHIP

```


1

 0 TIME STEP NUMBER 5
 TIME STEP BEGINNING JAN 5, 1983
 HMAX = -0.245137
 HMAX = -0.000079

NODE=	1	SPEC HEAD IN-OUT	0.00187
NODE=	2	SPEC HEAD IN-OUT	0.00375
NODE=	3	SPEC HEAD IN-OUT	0.00302
NODE=	4	SPEC HEAD IN-OUT	0.00144
NODE=	5	SPEC HEAD IN-OUT	-0.00164
NODE=	6	SPEC HEAD IN-OUT	-0.00704
NODE=	7	SPEC HEAD IN-OUT	-0.01519
NODE=	8	SPEC HEAD IN-OUT	-0.02471
NODE=	9	SPEC HEAD IN-OUT	-0.03035
NODE=	10	SPEC HEAD IN-OUT	-0.02473
NODE=	11	SPEC HEAD IN-OUT	-0.01521
NODE=	12	SPEC HEAD IN-OUT	-0.00705
NODE=	13	SPEC HEAD IN-OUT	-0.00164
NODE=	14	SPEC HEAD IN-OUT	0.00145
NODE=	15	SPEC HEAD IN-OUT	0.00305
NODE=	16	SPEC HEAD IN-OUT	0.00382
NODE=	17	SPEC HEAD IN-OUT	0.00417
NODE=	18	SPEC HEAD IN-OUT	0.00431
NODE=	19	SPEC HEAD IN-OUT	0.00435
NODE=	20	SPEC HEAD IN-OUT	0.00229

0 RIVER-NODE FLUXES

1	0.02524	33	0.00000
2	0.02482	34	0.00000
3	0.02620	35	0.00000
4	0.02990	36	0.00000
5	0.03732	37	0.00000
6	0.05011	38	0.00000
7	0.06797	39	0.00000
8	0.08227	40	0.00000
9	0.06789	41	0.00000
10	0.04996	42	0.00000
11	0.03709	43	0.00000
12	0.02955	44	0.00000
13	0.02562	45	0.00000
14	0.02374	46	0.00000
15	0.02295	47	0.00000
16	0.02278	48	0.00000
17	0.02309	49	0.00000
18	0.02420	50	0.00000
19	0.00000	51	0.00000

BUDGET BALANCE

	FT**3/S	FT**3
SPEC. HEAD INFLOW	0.0335	1.3646E+04

SPEC. HEAD OUTFLOW -0.1276 -2.4655E+04
 SPEC. BDRY INFLOW 0.2050 8.6400E+04
 SPEC. BDRY OUTFLOW 0.0000 0.0000E+01
 STORAGE CHANGE 0.4144 1.8902E+05
 PUMP INJECTN WELLS 1.0000 4.3200E+05
 PUMP DISCHRG WELLS -1.5474 -6.6843E+05
 APPLIED WATER 0.0000 0.0000E+01
 DIRECT RECHARGE 0.0000 0.0000E+01
 PRECIPITATION 4.8297 2.0844E+06
 EVAPOTRANSPIRATION -2.8793 -1.2439E+06
 GROUND-WATER E.T. -2.5940 -1.1271E+06
 1ST REACH LOSS 0.6707 2.5646E+05

 MASS RESIDUAL -5.4130E-05 -8.9200E+00

RC#		NODE		HEAD VALUES																			
1	1	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00		
200.00	2	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00		
200.22	2	199.97	199.98	199.98	200.00	200.00	200.00	200.02	200.07	200.14	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22		
200.14	3	200.07	200.02	200.00	199.98	199.98	200.00	200.06	200.16	200.33	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55		
200.33	4	200.16	200.06	200.00	199.97	199.96	200.01	200.10	200.26	200.56	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04		
201.04	4	199.95	199.96	199.98	200.01	200.10	200.26	200.56	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04		
200.56	5	200.26	200.10	200.01	199.97	199.96	200.01	200.10	200.26	200.56	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04	201.60	201.04		
201.60	5	199.95	199.96	199.98	200.02	200.12	200.32	200.73	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60		
200.73	6	200.32	200.12	200.02	199.98	199.96	200.02	200.12	200.32	200.73	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60		
201.04	6	199.95	199.96	199.98	200.02	200.10	200.26	200.56	201.04	201.59	201.04	201.59	201.04	201.59	201.04	201.59	201.04	201.59	201.04	201.59	201.04		
200.56	7	200.26	200.10	200.02	199.98	199.96	200.02	200.12	200.32	200.73	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60	203.72	201.60		
200.55	7	199.96	199.97	199.98	200.01	200.06	200.16	200.33	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55		
200.33	8	200.16	200.06	200.01	199.98	199.97	200.06	200.16	200.33	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71	200.55	200.71		
200.22	8	199.98	199.98	199.99	200.00	200.03	200.08	200.14	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22		
200.14	9	200.08	200.03	200.01	199.99	199.99	200.03	200.08	200.14	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27	200.22	200.27		
200.00	9	200.01	200.01	200.01	200.02	200.02	200.02	200.01	200.01	200.01	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00		
200.01	10	200.01	200.02	200.02	200.02	200.02	200.02	200.01	200.01	200.01	200.01	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00	200.00		
199.79	10	200.05	200.06	200.06	200.05	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03	200.03		
199.89	11	199.98	200.03	200.06	200.07	200.08	200.08	200.08	200.07	200.06	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05		
199.51	11	200.08	200.13	200.14	200.12	200.06	199.96	199.77	199.51	199.31	199.51	199.31	199.51	199.31	199.51	199.31	199.51	199.31	199.51	199.31	199.51		
199.77	12	199.96	200.06	200.12	200.15	200.16	200.16	200.16	200.14	200.09	200.09	200.09	200.09	200.09	200.09	200.09	200.09	200.09	200.09	200.09	200.09		
198.52	12	200.04	200.04	200.03	199.99	199.88	199.66	199.23	198.52	197.68	198.52	197.68	198.52	197.68	198.52	197.68	198.52	197.68	198.52	197.68	198.52		
199.23	13	199.66	199.88	199.99	200.04	200.07	200.07	200.07	200.06	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05	200.05		
199.51	13	199.99	199.99	199.97	199.90	199.77	199.47	198.86	197.53	194.25	197.53	194.25	197.53	194.25	197.53	194.25	197.53	194.25	197.53	194.25	197.53		
198.85	14	199.47	199.77	199.91	199.97	200.00	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01	200.01		
199.04	14	199.96	199.95	199.93	199.87	199.75	199.50	199.05	198.31	197.45	198.31	197.45	198.31	197.45	198.31	197.45	198.31	197.45	198.31	197.45	198.31		
199.04	15	199.50	199.75	199.88	199.94	199.96	199.97	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98	199.98		
199.36	15	199.95	199.94	199.92	199.88	199.79	199.62	199.36	199.01	198.76	199.01	198.76	199.01	198.76	199.01	198.76	199.01	198.76	199.01	198.76	199.01		
199.36	16	199.62	199.79	199.88	199.93	199.95	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96		
199.62	16	199.95	199.95	199.93	199.90	199.85	199.75	199.62	199.36	199.01	198.76	199.01	198.76	199.01	198.76	199.01	198.76	199.01	198.76	199.01	198.76		
199.62	17	199.75	199.85	199.90	199.94	199.95	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96	199.96		
199.62	17	199.98	199.97	199.96	199.94	199.91	199.86	199.79	199.73	199.70	199.73	199.70	199.73	199.70	199.73	199.70	199.73	199.70	199.73	199.70	199.73		
199.79	18	199.86	199.91	199.94	199.96	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97	199.97		
199.79	18	200.01	200.01	200.00	199.99	199.97	199.94	199.91	199.89	199.87	199.89	199.87	199.89	199.87	199.89	199.87	199.89	199.87	199.89	199.87	199.89		
199.91	19	199.94	199.97	199.99	200.00	200.00	200.01	200.01	200.02	200.02	200.02	200.02	200.02	200.02	200.02	200.02	200.02	200.02	200.02	200.02	200.02		
199.91	19	200.09	200.07	200.06	200.05	200.04	200.02	200.01	200.00	199.99	200.00	199.99	200.00	199.99	200.00	199.99	200.00	199.99	200.00	199.99	200.00		
200.01	20	200.02	200.04	200.05	200.05	200.06	200.06	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07	200.07		
200.38	20	200.20	200.16	200.14	200.13	200.12	200.11	200.10	200.09	200.08	200.09	200.08	200.09	200.08	200.09	200.08	200.09	200.08	200.09	200.08	200.09		
200.09	20	200.11	200.11	200.12	200.13	200.13	200.14	200.15	200.17	200.21	200.17	200.21	200.17	200.21	200.17	200.21	200.17	200.21	200.17	200.21	200.17		
200.09	20	200.11	200.11	200.12	200.13	200.13	200.14	200.15	200.17	200.21	200.17	200.21	200.17	200.21	200.17	200.21	200.17	200.21	200.17	200.21	200.17		

DRAWDOWNS FROM INITIAL HEAD VALUES

0 ROW NODE

