

A DUAL-POROSITY MODEL FOR SIMULATING  
SOLUTE TRANSPORT IN OIL SHALE

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## CONVERSION FACTORS

For those readers interested in using the metric system, the following table may be used to convert the inch-pound units of measurement used in this report to metric units:

<i>Multiply</i>	<i>By</i>	<i>To obtain</i>
foot	0.3048	meter
foot per day	0.3048	meter per day
foot squared per day	0.09290	meter squared per day

# A DUAL-POROSITY MODEL FOR SIMULATING SOLUTE TRANSPORT IN OIL SHALE

By Kent C. Glover

## ABSTRACT

A model is described for simulating three-dimensional ground-water flow and solute transport in oil shale and associated hydrogeologic units. The model treats oil shale as a dual-porosity medium by simulating flow and transport within fractures using the finite-element method. Diffusion of solute between fractures and the essentially static water of the shale matrix is simulated by including an analytical solution that acts as a source-sink term to the differential equation of solute transport. While knowledge of fracture orientation and spacing is needed to effectively use the model, it is not necessary to map the locations of individual fractures.

The computer program listed in the report incorporates many of the features of previous dual-porosity models while retaining a practical approach to solving field problems. As a result the report does not extend the theory of solute transport in any appreciable way. The emphasis in the report is on bringing together various aspects of solute-transport theory in a manner that is particularly suited to the unusual ground-water flow and solute-transport characteristics of oil-shale systems.

## INTRODUCTION

Digital models of ground-water flow and solute transport have been used to predict possible impacts of oil-shale development on the ground-water resource (Robson and Saulnier, 1981). Unfortunately, modeling techniques used in the past either were not developed for use in rock, such as oil shale, where porosity depends on fracturing, or required data that usually are not available. From 1980 to 1983 the U.S. Geological Survey studied migration of solute from an in situ oil-shale retort near Rock Springs, Wyo. (fig. 1). The emphasis of the study was to identify geologic, hydraulic, and chemical factors that control the process of solute transport in oil shale. During the study, it became apparent that no existing model was well suited for the unusual flow and transport characteristics of oil-shale systems. Therefore, as part of the study, a model was developed.

This report describes two digital-computer programs that are used to simulate ground-water flow and solute transport in oil shale with interbedded tuff or sandstone. The U.S. Department of Energy experimental, in situ, oil-shale retort near Rock Springs, Wyo., was used as the prototype flow system of the model. The application of modeling techniques to this flow system is described by Glover (1986). The model is designed for use in site-specific studies and consequentially includes many features, such as diffusion of solute between fractures and the shale matrix, that may not be important to basin-wide or regional investigations.

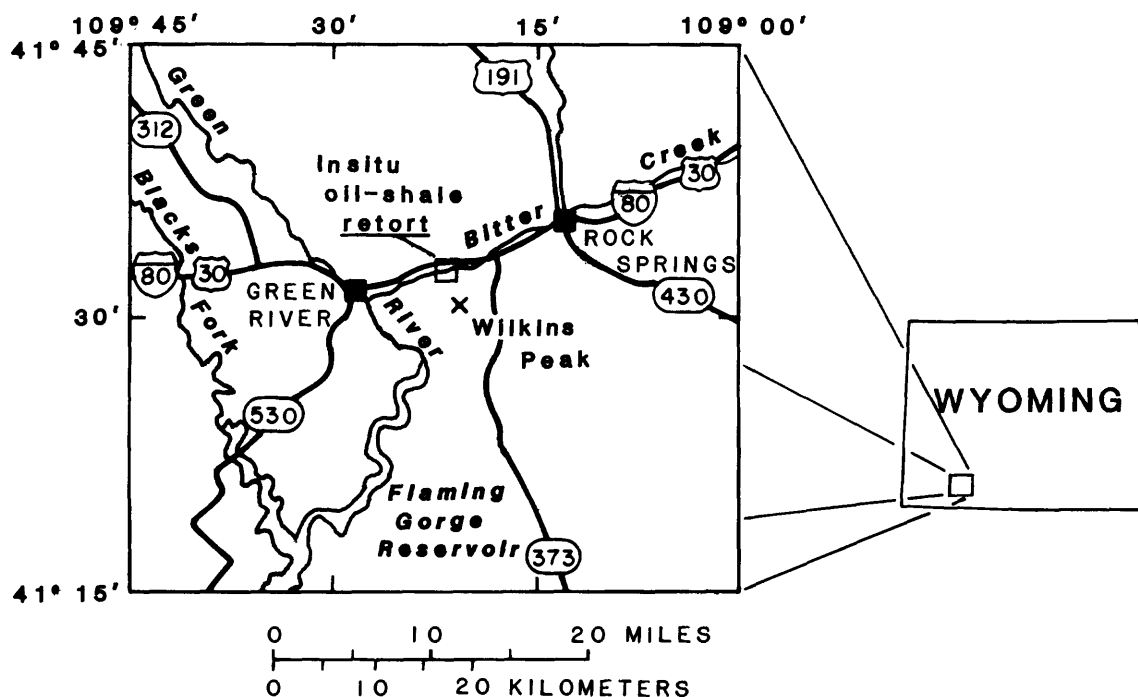


Figure 1.--Location of in situ oil-shale retort near Rock Springs, Wyoming.

A wide array of ground-water flow and solute-transport models have become available in recent years. Konikow and Bredehoeft (1978), and INTERCOMP Resource Development and Engineering, Inc. (1976) describe two of the more commonly used programs. Other models that are potentially useful in studies of oil-shale hydrology are described by Bibby (1981), Noorishad and Mehran (1982), and Rasmuson and others (1982). The model described in this report incorporates many of the features of the above models while retaining a practical approach to the solution of field problems. Boundary conditions can be applied easily and mapping of individual fractures is not necessary. As a result of this approach, the present model does not extend the theory of solute transport in fractured media in any appreciable way. Instead, various aspects of this theory are brought together in a manner that is particularly suited to the unusual ground-water flow and solute-transport characteristics of oil-shale systems.

Throughout the remainder of this report it is assumed that the reader is familiar with the basics of solute-transport modeling in porous media, such as sand-and-gravel aquifers, as well as the finite-element method of numerical analysis. It is believed that most hydrologists faced with a problem of solute migration in oil shale will have faced similar problems in unfractured media. For those readers unfamiliar with applications of solute transport modeling using the finite-element method, it is suggested that teaching references such as Freeze and Cherry (1979), Konikow and Bredehoeft (1978), and Zienkiewicz (1971) be consulted.

#### OIL SHALE AS A DUAL-POROSITY MEDIUM

Porosity in oil-shale formations may be classified on the basis of its relationship to hydraulic conductivity. Porosity may be considered effective for increasing hydraulic conductivity if the pores are interconnected or not effective, as in shale, if the pores are relatively isolated. Porosity due to faults, joints, collapse breccia, and solution cavities is the major source of permeability in oil shale with otherwise low permeability (Robson and Saulnier, 1981). Persistent layers of permeable tuff or sandstone, ranging in thickness from less than an inch to several inches, also are common in oil shale (Bradley, 1964) and contribute to the hydraulic conductivity. The porosity of the shale and marlstone matrix, although relatively high, does not contribute significantly to hydraulic conductivity.

Any attempt to simulate ground-water flow and solute transport in oil shale must include consideration of the dual-porosity nature of these sediments. The solute-transport processes of hydrodynamic dispersion and advection are related directly to seepage velocity and therefore are related to the effective porosity of the formation. However, several investigators including Grisak and Pickens (1980) and Bibby (1981) have noted that dispersion and advection alone cannot account for the distribution of non-reacting solute in fractured formations. The extremely low seepage velocity of water within the shale matrix minimizes dispersion and advection, increasing the relative importance of diffusion. The porosity of the shale matrix, although not contributing to hydraulic conductivity, is important to dispersion.

## METHODS OF SIMULATION

### Ground-Water Flow

Only the part of formation porosity that contributes to hydraulic conductivity needs to be considered in analysing ground-water flow systems. For most practical purposes this porosity in oil shale, whether resulting from faulting, solution channels, or thin beds of permeable tuff and sandstone, can be analysed using the techniques of Snow (1969). He showed that many problems of flow through dual-porosity media can be solved using an anisotropic hydraulic-conductivity tensor in conjunction with standard porous-media techniques. His approach to the problem of ground-water flow is valid if the formation has a fracture density that is high compared to the scale of the problem. In such a case, the hydraulic characteristics of the fractured formation are similar to those of granular media. If fracture spacing is irregular in a given direction, the formation will exhibit heterogeneity. Because mapping of individual fractures is impractical for most field problems, the approach of Snow (1969) is used in this model. From a practical viewpoint the model described in this report may be useful when the hydrologist has some knowledge of the average spacing, aperture size, and directions of fracturing, but accurate mapping of individual fractures is not possible. Through proper alignment of the model cartesian coordinate axes with principal directions of fracturing and use of heterogeneous hydraulic conductivity values, many different fracture geometries can be simulated.

The three-dimensional nature of a fracture system, along with the essentially horizontal bedding plane of most formations, generally results in a three-dimensional flow system in oil-shale strata. Three-dimensional flow systems in oil shale have been observed on both a regional scale by Robson and Saulnier (1981) and on a local scale by the author at the Department of Energy in situ oil-shale retort near Rock Springs, Wyo. Therefore, the program given in this report is intended for use in studies of three-dimensional ground-water flow.

Rocks in the vicinity of an in situ retort are altered due to induced fracturing. As a result fracture patterns in the retort chamber are significantly different than regional fracture patterns. The values and degree of anisotropy of hydraulic conductivity within a retort chamber also differ from regional estimates. Nevertheless the fracture density usually is high, compared to the dimensions of the retort chamber, and standard porous-media techniques can be used to simulate ground-water flow in the immediate vicinity of the retort chamber.

### Solute Transport

The extension of solute-transport theory to dual-porosity media has attracted considerable attention in water-resources literature. Much of the literature has been directed toward studies of radionuclide migration through individual fractures and as such is not directly applicable to field problems where mapping of individual fractures is impractical. Grisak and Pickens (1980) modeled fractured media by using separate finite elements to represent



fractures and adjacent low permeable rock. Finite elements used to represent the fractures were modeled with material properties that differed from adjacent low-permeability elements. This approach was extended by Noorishad and Mehran (1982), who also introduced the use of an "upstream-weighting" technique to simulate more accurately the essentially advective transport that occurs within the fractures.

Bibby (1981) took a different approach to modeling solute movement through dual-porosity media. Molecular diffusion between fractures and the primarily static water in the shale matrix was incorporated in the model as a source-sink term utilizing an appropriate analytical solution. To apply this two-dimensional model the user needs to know average spacing, aperture, and direction of fractures but does not need to map individual fractures. This approach is analogous to the method of flow analysis discussed by Snow (1969). Although no upstream-weighting technique was used, Bibby (1981) reported no difficulty in applying the model to a field problem in a limestone aquifer.

Rasmuson and others (1982) presented a model for simulating three-dimensional solute transport in fractured rocks that is based on an integrated finite-difference approach. The model produces excellent results if solute transport is dominated by hydrodynamic dispersion. Because advection dominates most solute-transport problems in oil-shale strata (Robson and Saulnier, 1981), the approach of Rasmuson and others (1982) is not considered practical for use with oil shale.

The approach to modeling solute transport presented in this report is basically the model described by Bibby (1981) and extended to three dimensions. This approach has been selected because it strikes a balance between theoretical accuracy and practical applicability. Accurate mapping of fractures is not needed to use the model, although some knowledge of the fracture system is required. The ease with which boundary conditions can be applied also supports this approach. One disadvantage of the approach is the difficulty of extending the analytical solution for diffusion between fractures and the shale matrix to problems of multiple-species transport.

A model that can be extended to problems of multiple-species transport is described by Huyakorn and others (1983). Unfortunately, fracture locations must be accurately mapped to use the model successfully. Therefore the model of Huyakorn and others (1983), although theoretically superior to the model of Bibby (1981), cannot be applied in most practical field problems.

## APPLICATION OF THE FINITE-ELEMENT METHOD

### Ground-Water Flow

The basic governing equation for three-dimensional flow, when Cartesian coordinate axes are aligned with the principal components of the hydraulic-conductivity tensor, is as follows:

$$\frac{\partial}{\partial x_i} (K_i \frac{\partial h}{\partial x_i}) + W = S_s \frac{\partial h}{\partial t} \quad i = 1, 2, 3 \quad (1)$$

where  $K_i$  = the hydraulic-conductivity tensor [ $LT^{-1}$ ];  
 $W$  = the source-sink function (positive for a source) [ $T^{-1}$ ];  
 $S_s$  = specific storage [ $L^{-1}$ ];  
 $h$  = the hydraulic head [ $L$ ];  
 $x_i$  = Cartesian coordinate [ $L$ ], and use of a repeated subscript indicates summation on that subscript; and  
 $t$  = time [ $T$ ].

The source-sink term may be distributed areally or may represent a well. Boundary conditions that may be applied on the periphery of the problem area include known specific discharge normal to the boundary or known hydraulic head. Parameters  $K_i$ ,  $W$  and  $S_s$  are approximated by subdividing the region of interest into discrete zones. Parameters are assumed to be constant within each zone which gives rise to internal boundary conditions at zonal discontinuities. Along these internal boundaries both hydraulic head and normal specific discharge must remain unchanged as the boundary is crossed.

Equation 1 may be solved by using the finite-element method based on the Galerkin criteria. Details of this method are described in Zienkiewicz (1971) and Pinder and Gray (1977). Within this report, first-order basis functions are used to describe each cubic element. While use of these basis functions requires more elements than if higher-order functions were used, the reduced oscillatory behavior and integration time associated with linear functions make this simpler approach preferable.

The matrix equation resulting from the finite-element discretization of equation 1 is as follows:

$$\left( \frac{1}{\Delta t} \bar{\bar{C}} + \theta \bar{\bar{K}} \right) \bar{h}^m = \bar{q} + \left( \frac{1}{\Delta t} \bar{\bar{C}} - (1-\theta) \bar{\bar{K}} \right) \bar{h}^{m-1} \quad (2)$$

where  $\bar{\bar{C}}$  = a coefficient matrix involving specific storage ( $S_s$ ) and calculated for a single finite element as  $\iiint_V S_s \bar{n} (\bar{n})^t dV$ ;  
 $\bar{\bar{K}}$  = a coefficient matrix, with off-diagonal components equal to zero, involving hydraulic conductivity and calculated for a single finite element as

$$\iiint_V (K_i \frac{\partial \bar{n}}{\partial x_i} (\frac{\partial \bar{n}}{\partial x_i})^t) dV, \quad i = 1, 2, 3;$$

$\bar{h}^m$  = the vector of hydraulic head at time m;

$\bar{q}$  = the known vector involving source-sink terms and specified flux boundary conditions;

$\partial t_m$  = the length of the time step;

$\theta$  = a number ranging from 0 for an explicit solution to 1 for an implicit solution;

$V$  = the volume of the finite element;

$\bar{n}$  = the vector of finite-element shape functions; and

$( )^t$  = the transpose of the enclosed vector.

Equation 2 is solved by Gaussian elimination for banded symmetric matrices. Although other direct-solution techniques, such as Cholesky decomposition, can be used, the procedure used in this report has proven adequate. Gupta and Tanji (1976) has suggested the use of matrix solvers that take full advantage of the sparse nature of coefficient matrices in attempts to reduce core-storage requirements. The tradeoffs in these matrix solvers are increased disk access and increased computational time. With the introduction of virtual-memory operating systems used by most computers, such detailed attention to minimizing core storage is not required.

### Solute Transport

The transport of a conservative solute in ground water is described by the following equation:

$$\frac{\partial}{\partial x_i} (D_{ij} \frac{\partial c}{\partial x_j}) - \frac{\partial}{\partial x_i} (cq_i) = \phi \frac{\partial c}{\partial t} - Wc^* \quad i, j = 1, 2, 3 \quad (3)$$

where  $D_{ij}$  = the hydrodynamic-dispersion coefficient [ $L^2T^{-1}$ ];

$c$  = solute concentration [ $ML^{-3}$ ]; and

$\phi$  = porosity [dimensionless];

$q_i = -K_i \frac{\partial h}{\partial x_i}$  = the Darcian fluid velocity [ $LT^{-1}$ ];

$W$  = a source-sink function [ $T^{-1}$ ];

$c^*$  = solute concentration in a fluid source [ $ML^{-3}$ ].

When applied to dual-porosity media such as oil shale, the terms are defined relative to the joint or fracture system, with the shale matrix considered impermeable. The hydrodynamic-dispersion coefficient ( $D_{ij}$ ) is related to Darcian fluid velocity in the fracture system by

$$D_{ij} = a_{ijmn} \frac{q_m q_n}{|q|} + D^* \quad i, j, m, n, = 1, 2, 3$$

where  $a_{ijmn}$  = dispersivity of the fractured media [L],

$q_m$  and  $q_n$  = Darcian fluid velocity [ $LT^{-1}$ ],

$|q|$  = the magnitude of the Darcian velocity vector [ $LT^{-1}$ ], and

$D^*$  = the molecular-diffusion coefficient [ $L^2T^{-1}$ ].

for isotropic media,

$a_{iiii} = \alpha_L$  = the longitudinal dispersivity for  $i = 1, 2, 3$ ;

$a_{ijij} = \alpha_T$  = the transverse dispersivity for  $i, j = 1, 2, 3$ ; and

$a_{ijji} = \frac{1}{2} (\alpha_L - \alpha_T)$  for  $i, j = 1, 2, 3$ .

There is no corresponding theory describing hydrodynamic dispersion in anisotropic media.

The form of the transport equation used in this report also is used by Pinder and Gray (1977), and Bibby (1981). Konikow and Bredehoeft (1978) and most studies of solute transport done by the U.S. Geological Survey use a form of the equation that is obtained by dividing equation 3 by porosity. The hydrodynamic-dispersion coefficient then is defined in terms of seepage velocity. Either form of the equation can be used successfully. However, when comparing results of studies using different forms of the transport equation, it should be recognized that the dispersion coefficients will differ by the magnitude of porosity.

Galerkin's method of weighted residuals, when applied to equation 3 with isoparametric finite elements, gives

$$\left(\frac{1}{\Delta t_m} \bar{P} + \theta \bar{D}\right) \bar{c}^m = q_c^* + \left(\frac{1}{\Delta t_m} \bar{P} - (1-\theta)\bar{D}\right) \bar{c}^{m-1} \quad (4)$$

where  $\bar{P}$  = a coefficient matrix involving porosity and is calculated for a single finite element as  $\iiint_V \phi \bar{n} (\bar{n})^t dV$ ;

$\bar{D}$  = a coefficient matrix involving hydrodynamic dispersion and advection and is calculated for a single finite element as

$$\iiint_V (D_{ij} \frac{\partial \bar{n}}{\partial x_i} (\frac{\partial \bar{n}}{\partial x_j})^t - q_i \frac{\partial \bar{n}}{\partial x_i} (\bar{n})^t) dV \quad i, j, = 1, 2, 3;$$

$\bar{c}^m$  = the vector of solute concentration in fractures at time  $m$ ; and

$q_c^*$  = a known vector involving specified flux boundary conditions and source-sink terms ( $W$ ), solute concentration in fluid sources ( $c^*$ ), and diffusion between fractures and the rock.

Detailed information on formulating this equation can be obtained from Bibby (1981) or Pinder and Gray (1977, p. 144-148). Equation 4 is solved by the Gauss-Doolittle method for banded, nonsymmetric matrices.

#### Diffusion Within the Oil-Shale Matrix

The exchange of solute between water in a system of parallel fractures and the essentially static water in the adjacent oil-shale matrix can occur by molecular diffusion and can be simulated by the addition of a source-sink term to equation 4. Bibby (1981) gives this source-sink term as a convolution integral that is expressed here in discrete form at time  $t_m$ .

$$W_c^* = - \frac{\phi_B b}{\Delta t_m (f+b)} [c_B(t_m) - c_B(t_{m-1})] \quad (5)$$

where  $W_c^*$  = the mass flux of solute entering the fracture  $[M/L^3/T]$ ;

$\phi_B$  = the porosity of the oil-shale matrix [dimensionless];

$b$  = the average distance between parallel fractures  $[L]$ ;

$\Delta t_m$  =  $[t_m - t_{m-1}]$ , the length of the time step  $[T]$ ;

$f$  = the average fracture width  $[L]$ ;

$c_B$  = the average solute concentration in the oil-shale matrix  $[M/L^3]$ ;

$$\begin{aligned}
c_B(t_m) = & c(t_0) \left[ 1 - \frac{8}{\pi^2} \sum_{n=0}^{\infty} \frac{1}{E_n} \exp(-B_n t_m) \right] + \sum_{i=1}^m \{ c(t_i) \\
& - c(t_{i-1}) \} \left\{ 1 - \frac{8}{\pi^2 \Delta t_i} \sum_{n=0}^{\infty} \frac{1}{E_n B_n} \{ \exp[-B_n(t_m - t_i)] \right. \\
& \left. - \exp[-B_n(t_m - t_{i-1})] \} \right\} ;
\end{aligned}$$

$c(t_0)$  = the solute concentration in the fracture at the beginning of the simulation  $[M/L^3]$ ;

$n$  = an index of summation [dimensionless];

$$E_n = (2n - 1)^2;$$

$$B_n = D_d(2n+1) \pi^2/b^2;$$

$c(t_i)$  = the solute concentration in the fracture at time  $t_i$   $[M/L^3]$ ; and

$D_d$  = molecular diffusion of solute within the oil-shale matrix  $[L^2/T]$ .

The only unknown variable in equation 5 is  $c(t)$ . Therefore, the coefficient associated with  $c(t)$  is placed on the main diagonal of the  $\bar{D}$  matrix in equation 4, while the remainder of the expression is placed in the righthand vector of known values. Equation 5 involves summation over the entire period of the simulation. Fortunately, little truncation error is introduced by retaining only a small number of terms in the summations. The optimum number of terms to retain depends on the values of  $D_d$ ,  $b$  and  $\Delta t$ ; however, from information presented by Bibby (1981), it is apparent that between five and ten terms are adequate for most applications.

Several aquifer properties must be determined to use equation 5 successfully. These properties include molecular diffusion within the oil-shale matrix ( $D_d$ ), average distance between parallel fractures ( $b$ ), and average fracture width ( $f$ ). Because direct estimates of these properties by field tests usually are unreliable, model calibration must be used to improve upon the estimates. Therefore equation 5 generally is useful only in studies with well distributed and frequent measurements of solute concentration in wells.

## Sources of Solute Within the Retort Chamber

Mechanisms for creating and leaching chemicals from burned shale within an in situ retort chamber to the formation water are poorly understood. Knowledge of conditions during the retort process is needed to determine the chemicals that are created, while an understanding of dissolution controls is needed to predict how the created chemicals are transferred from the burned shale to the formation water.

The retort temperature, total time of retorting, permeability of the retort chamber, and initial mineralogy affect the type and amount of chemicals created during a burn. During combustion of a retort chamber, a "flame front" or zone of combustion moves through the fractured oil-shale bed. Hot gases from this combustion move ahead of the flame front and provide energy for pyrolysis of kerogen. In the process, a number of chemical byproducts are created. A discussion of how these byproducts are created is outside the scope of this report. However, it is important to recognize that the resulting chemicals will vary from retort chamber to retort chamber. In large chambers, the distribution of chemical byproducts within the chamber also may be important. Coring of an abandoned retort chamber before it resaturates and lab analysis of the cores by long-term leaching studies can provide some understanding of the chemicals that are created during retorting.

After the retort operation ceases, the chamber resaturates and chemicals created during the burn are leached into the water. During the time that it takes to resaturate the retort chamber a rapid increase in the concentration of solute occurs. For many studies of solute transport, an equation to describe the initial dissolution is not needed because migration of solute from the retort chamber is unlikely until the formation resaturates.

After the retort chamber resaturates, as Hall (1982) proposed, diffusion between the pores of the oil-shale matrix and fractures acts as a dominant mechanism for introducing solute to a ground-water system. This mechanism is simulated by equation 5. As discussed previously, equation 5 describes diffusion between the oil-shale matrix and a system of parallel fractures. Fracturing within a retort chamber is much more irregular but can be conceptualized as two systems of parallel fractures, one vertical and one horizontal, resulting in cubes of oil shale. Matrix diffusion within the retort chamber can be simulated by applying equation 5 once for each fracture orientation.

In addition to rock and fracture characteristics discussed previously, the use of equation 5 to simulate a source of solute requires an estimate for the initial solute concentration in the pores of the oil-shale matrix ( $c_B$ ). Water-quality sampling as the retort chamber resaturates will provide a measure of solute concentration in fractures but not of solute concentration in the shale-matrix pores. In cases where solute migration outside the chamber does not occur until some time after the initial resaturation the initial solute concentration within the shale-matrix pores may be approximated by the initial solute concentration within the fractures. Study by the author and the Department of Energy of an in situ oil-shale retort near Rock Springs, Wyo., indicate that this approximation may be reasonable in many cases.

In situ retort chambers are sources of solute for extended periods of time after retorting stops. This characteristic has been observed by the author at the U.S. Department of Energy experimental in situ oil-shale retort near Rock Springs, Wyo., and by Hall (1982) in laboratory experiments. The source of solute cannot be explained solely by diffusion from burned shale into fractures.

Hall (1982) has proposed a mechanism to explain the long-term source of solute that is based on the slow dissolution of the mineral matrix from pore walls. Although he was unable to provide a good theoretical basis for this slow dissolution, he proposed the use of an empirical expression to simulate the mechanism. The expression is

$$K^* A (c_e - c) \quad (6)$$

where  $K^*$  = the mass-transfer coefficient [ $L^2 T^{-1}$ ],  
 $A$  = the interface area [ $L^2$ ],  
 $c_e$  = the equilibrium concentration of fluid in contact  
with shale blocks [ $ML^{-3}$ ], and  
 $c$  = the solute concentration [ $ML^{-3}$ ].

This equation acts as a source-sink term to the solute-transport equation and simulates the dominant mechanism for leaching of burned shale after the passage of 2 to 3 pore volumes of water.

Application of the proposed leaching mechanism to a field problem requires that  $K^*$ ,  $A$ , and  $c_e$  in equation 6 be determined. While Hall (1982) determined values of the product  $K^*A$  for organic solute from column experiments, the applicability of these values to field problems is unknown. Therefore  $K^*A$  is effectively a parameter that must be estimated during model calibration. The equilibrium concentration,  $c_e$ , can be estimated from long-term leaching tests of burned shale. Unfortunately, these tests have been performed for very few of the chemical species found in oil-shale retort water.

The use of empirical equation 6 to describe mass transfer essentially provides an infinite source of solute to the medium where, in reality, such a source is finite. Therefore the use of equation 6 will overestimate solute concentrations in the trailing part of a plume. The use of equation 6 is justified in studies of solute transport in oil shale where the distribution of solute in the trailing part of the plume is not considered, or in studies where an estimate of maximum probable concentration of solute in the trailing part of a plume is needed.



## MODEL CALIBRATION AND SENSITIVITY

The approach to ground-water flow and solute transport used in this report along with the three-dimensional nature of flow through oil shale, introduces an unusually large number of calibration parameters. Considering the amount and distribution of water-level and solute-concentration data that are available in most field studies, it is possible to reach a point where additional detail in the simulation procedure does not significantly improve the model fit. Therefore, it is important to assess the uncertainty associated with the various calibration parameters if the model is to remain a practical one for use in field problems.

Cooley (1977) presents a method for evaluating the reliability of a model within the framework of steady-state flow of ground water in two dimensions. This technique has been extended to three dimensions and is described in a later section. Before presenting the technique, the relation between measured data in wells and model-calculated hydraulic head is discussed. Observation wells that have been installed in oil-shale strata rarely are piezometers.

### Relation of Measured Data and Calculated Results

#### Hydraulic Head

The relation between hydraulic head in the formation and the water level measured in a well open to part or all of the formation is governed by aquifer properties, well-bore characteristics and the vertical-head gradient within the formation. An accurate treatment of the relation would involve solving the three-dimensional equation of ground-water flow in the close vicinity of the well bore using model estimates of aquifer properties and a boundary condition of uniform head along the well bore. The computed head in the well could be compared to measured water-level data. This approach is not practical in most field problems. Instead, a relation is used in the model that is based on the steady-state conservation of water within the well bore and Darcy's law. The result is a simple weighted average of hydraulic head.

$$h_{int} = \frac{\int_{z_1}^{z_2} K_{ii} h dz}{\int_{z_1}^{z_2} K_{ii} dz} \quad i = 1, 2 \quad (7)$$

where  $h_{int}$  = the depth-integrated head in the well [L],

$h$  = hydraulic head obtained from the solution to equation 2 [L],

$z_2$  = the altitude of the top of the well bore open to the formation [L],

$z_1$  = the altitude of the bottom of the well bore open to the formation [L], and  
 $K_{ii}$  = hydraulic conductivity [ $LT^{-1}$ ].

Equation 7 is derived by writing Darcy's law for radial flow and integrating along the open interval of the well. Applying Darcy's law for radial flow from a circular boundary with no drawdown due to well-bore effects and assuming a well radius of 1 foot gives

$$q = \frac{2\pi K(h_b - h_{int})}{\ln(r)}$$

where  $q$  = the Darcian flux per unit length of well bore [ $L^2T^{-1}$ ],  
 $K$  = the hydraulic conductivity [ $LT^{-1}$ ],  
 $h_b$  = the head at the boundary [L],  
 $h_{int}$  = the head in the well [L], and  
 $r$  = the distance to the boundary [L].

The distance  $r$  may be thought of as the radius of influence by the well bore. By assuming steady-state flow along the well bore, the amount of water entering the well along part of the open interval must equal the amount leaving along the remainder of the open section. The integral expression for this conservation of mass is

$$\int_{z_1}^{z_2} q \, dz = 0 .$$

Substitution of Darcy's law for radial flow into the expression for conservation of mass gives

$$\int_{z_1}^{z_2} \frac{K}{\ln(r)} h_b \, dz = h_{int} \int_{z_1}^{z_2} \frac{K}{\ln(r)} \, dz .$$

Treating distance to the boundary as a constant,  $\ln(r)$ , may be moved outside the integral and equation 7 obtained.

The model in this report uses information on the open intervals of wells, estimates of hydraulic conductivity, and calculated hydraulic head to determine the depth-integrated head at each observation well. The result may be compared to measured water-level data as a guide during model calibration when piezometers are not available.

## Solute Concentration

The relation between solute concentration in the formation and that in a well is governed by the three-dimensional equation of solute transport. As with hydraulic head, it is not practical to solve this problem at each well bore. By assuming steady-state conservation of solute and using Darcy's law, one can obtain an expression similar to equation 7.

$$c_{int} = \frac{\int_{z_3}^{z_4} c q_i dz}{\int_{z_3}^{z_4} q_i dz} \quad i = 1, 2 \quad (8)$$

where  $c_{int}$  = the depth-integrated solute concentration  $[M/L^3]$ ,  
 $c$  = solute concentration obtained from the solution to equation 4  $[M/L^3]$ ,  
 $z_3$  to  $z_4$  = the part of the well bore where water enters  $[L]$ , and  
 $q_i = K_{ii}(h - h_{int})$ .

Equation 7 is used in the model to infer sections of the well bore where water enters. The derivation of equation 8 essentially is identical to the derivation of equation 7 and will not be given here.

### Estimating Parameters for Models of Steady-State Ground-Water Flow

Direct measurement of all hydrogeologic parameters needed to construct models of ground-water flow and solute transport is rarely possible. As a result, values of unmeasured parameters usually are adjusted until measured and calculated water levels and solute concentrations match in some acceptable manner. Cooley (1977) proposed a regression method for estimating an optimal set of hydrogeologic parameters and assessing parameter reliability within the framework of two-dimensional steady-state flow. An extension of this method to three dimensions is included in the model described in this report.

The following method is used to estimate parameters and assess reliability for a three-dimensional model of steady-state ground-water flow. Throughout the following development, deviations from the two-dimensional development of Cooley (1977) are noted. The method is an iterative technique that minimizes the squared difference between measured and calculated water levels in wells. To set up the iterative technique it is necessary to linearize the finite-element form of the equation of ground-water flow with respect to unknown model parameters, differentiate the linearized equation with respect to these parameters, and, setting each derivative equal to zero, solve the system of equations.

The derivatives of unknown model parameters are obtained in the following manner. A truncated Taylor series expansion of equation 2 is written for  $(\Delta t)^{-1}$  equal to zero and  $\theta$  equal to one. The resulting iterative equation is identical to the one given by Cooley (1977). Solving this Taylor series for hydraulic head gives

$$\bar{h}^{r+1} = \bar{h}^r + \bar{\Delta}^r + \bar{S}^r \bar{b}^{r+1} \quad (9)$$

where  $\bar{h}^{r+1}$  = the vector of hydraulic head at the  $r+1$  iteration;

$\bar{h}^r$  = the vector of hydraulic head at the  $r$  iteration,

$\bar{\Delta}^r = (\bar{K}^r)^{-1} \bar{f}^r$ ;

$\bar{K}^r$  = the matrix involving estimates of hydraulic conductivity used in equation 2 at the  $r$  iteration;

$\bar{f}^r$  = the functional representative of equation 2 at the  $r$  iteration,

$\bar{S}^r = -(\bar{K}^r) \left( \frac{\partial \bar{f}^r}{\partial \bar{a}^r} \bar{a}^r \right)$ ;

$\bar{a}^r$  = the vector of model-parameter estimates at the  $r$  iteration,  
and

$\bar{b}^{r+1} = \frac{\bar{a}^{r+1} - \bar{a}^r}{\bar{a}^r}$ .

This equation is identical to the one used by Cooley (1977) to calculate hydraulic heads for each new iteration. The definitions of the variables are given in three dimensions instead of two.

Hydraulic head calculated by equation 9 cannot be compared directly to water-level measurements in wells unless the wells are constructed as piezometers. Because wells usually are open to some interval of aquifer, water levels represent a depth-integrated value of hydraulic head. Therefore, the least-squares criterion that must be satisfied differs slightly from the criterion used by Cooley (1977), where vertical-head variations were not considered.

$$\frac{\partial}{\partial \bar{b}^{r+1}} \{ \bar{w}(\bar{h}_m - \bar{h}_{int}^{r+1})^2 \} = 0 \quad (10)$$

where  $\bar{w}$  = a vector of weights between 0 and 1 describing the reliability of  $\bar{h}_m$ ,

$\bar{h}_m$  = the measured head, and

$\bar{h}_{int}^{r+1}$  = the vector of depth-integrated head at the  $r+1$  iteration.

Depth-integrated head is calculated by using the finite-element approximation to equation 7 in conjunction with equation 9. Substitution of the result in equation 10 gives the following:

$$\bar{w}(\bar{N}^r \bar{S}^r)^t \bar{N}^r \bar{S}^r \bar{b}^{r+1} = (\bar{N}^r \bar{S}^r)^t \bar{w}[\bar{h}_m - \bar{N}^r(\bar{h}^r + \bar{\Delta}^r)] \quad (11)$$

where  $\bar{N}^r$  = the matrix of finite-element approximations of equation 7 involving estimated hydraulic conductivity at the  $r$  iteration, and,

$( )^t$  = the matrix-transpose operation.

With the exception of the  $\bar{N}^r$  term, equation 11 is identical to the least-squares equation used by Cooley (1977). Equation 11 can be solved for  $\bar{b}^{r+1}$  which in turn can be used to calculate estimates of model parameters. Because the Taylor series expansion used to linearize the equations of ground-water flow is an approximation, the solution of equation 11 may not give global optimum values of  $\bar{b}^{r+1}$  unless  $\bar{h}^{r+1} - \bar{h}^r$  and  $\bar{a}^{r+1} - \bar{a}^r$  are small. Therefore iteration must be used.

The iteration procedure is identical to the one used by Cooley (1977). An initial estimate of aquifer properties and other unknown model parameters is used to solve equation 2 for hydraulic head. The coefficients  $\partial \bar{f}^r / \partial \bar{a}^r$  are calculated and equation 11 is used to solve for  $\bar{b}^{r+1}$  and new estimates of model parameters. The iteration number is advanced by one and the Taylor series expansion of equation 2 is used to compute  $\bar{f}^r$ . An iterative cycle is begun by resolving equation 2 with the latest estimates of aquifer properties and boundary flux rates until values for  $\bar{f}^{r+1}$  are less than some acceptable error.

## Sensitivity Analysis of Transient Ground-Water Flow and Solute Transport

Development of a statistical procedure for estimating model parameters under transient conditions is more difficult than under steady-state conditions. Several approaches to solving this problem may be possible, but all have proven overly expensive in terms of disk storage, core storage, or computational time. Therefore, calibration of transient ground-water flow and solute-transport models remains a trial-and-error process in this report.

A review of the differential equation of solute transport (eq. 3) shows that ground-water velocity acts as a calibration parameter. Therefore, it usually is necessary to develop a ground-water flow model concurrently with the development of a solute-transport model. Darcy's law is used to compute velocity from estimates of hydraulic conductivity and model-calculated head. The linkage between ground-water flow and solute transport often forces the hydrologist to iterate between the flow and transport models until a distribution of hydraulic conductivity is obtained that reproduces both historical water-level and solute-concentration data. Bibby (1981) observed that the solute-transport model can have a strong influence in determining optimal estimates of hydraulic conductivity.

### EVALUATION OF MODEL

A number of tests of the model have been made including comparisons of model results to analytical solutions and to one field application. Because analytical solutions rarely exist for three-dimensional problems, tests have been designed to verify the accuracy of the model in one and two dimensions. Repeated applications of the test problems in each of the three directions were used to isolate any programming or logic errors. Comparisons of model results with analytical solutions are discussed below. Application of the model to a field problem, solute transport from an in situ oil-shale retort near Rock Springs, Wyo., is discussed in Glover (1986).

### Ground-Water Flow

An evaluation of the basic model of ground-water flow (eq. 2) has not uncovered any unusual characteristics. For problems with relatively regular boundaries, the evaluation showed that the finite-element method did not improve upon head distribution calculated by finite-difference methods. In problems with complex and irregular aquifer geometry, the finite-element method can be used to model the flow system with fewer nodes.

Model results were compared to Theis' analytical solution for drawdown in the vicinity of a pumping well (Lohman, 1979). The Theis solution is for a fully penetrating well in an infinite confined aquifer with no vertical movement of water. These conditions were simulated in the flow model by using variable node spacing, a zero-drawdown boundary 20,000 ft from the well, and uniform aquifer properties. The aquifer was assumed to be 100 ft thick with a hydraulic conductivity of 12.5 ft/d and specific storage of  $10^{-6}$  ft<sup>-1</sup>. Well

discharge, distributed uniformly along the well bore, was simulated at a rate of 2.0 ft<sup>3</sup>/sec. Model results and the analytical solution are plotted in figure 2 for three distances from the pumping well.

The remainder of this section is a discussion of the regression technique for estimating aquifer properties and boundary-flow rates in steady-state systems. Most of the major conclusions of Cooley (1977 and 1979) seem to be appropriate for flow in three as well as two dimensions. Specific points of interest to three-dimensional problems are discussed below.

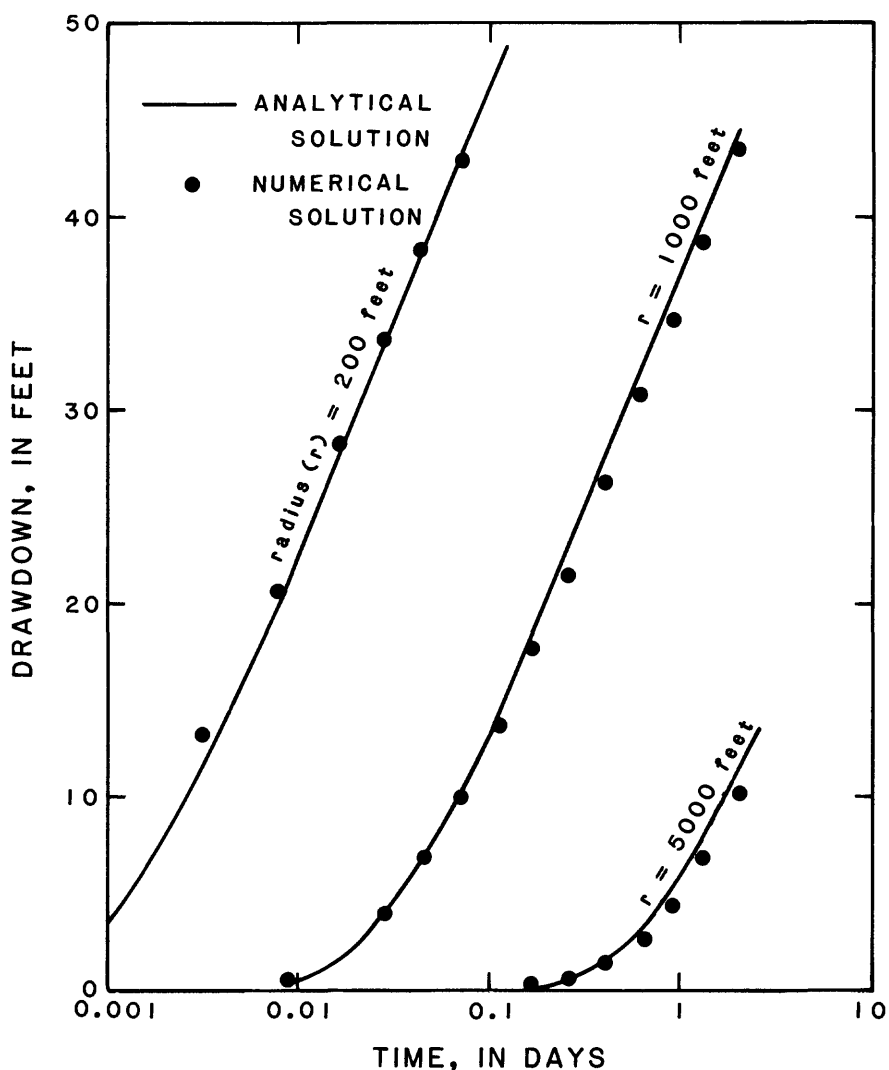


Figure 2.--Comparison of flow-model results with Theis' analytical solution.

The number of aquifer properties and flux rates that are treated as regression parameters can have a strong influence on the convergence characteristics of the model. The vertical anisotropy that is common in most three-dimensional systems increases the number of regression parameters and can cause convergence problems more easily than in two-dimensional flow. The number of regression parameters in three dimensions also increases because horizontal hydraulic conductivity often varies with depth. With the large number of parameters that can occur in three-dimensional models it may be difficult to obtain adequate water-level data for all strata and convergence can be very slow. In cases where the number of regression parameters approaches the number of measured water levels, no solution may be possible. These characteristics can limit the use of the regression technique in three-dimensional problems.

Boundary-flux rates usually are measured with greatest accuracy when the flow occurs at or near the land surface. This fact has a number of implications in using the regression technique. If significant recharge and discharge occur at depth as underflow and the hydraulic conductivity is not known exactly, the least-squares matrix usually is poorly conditioned and no solution is possible. On the other hand, if boundaries are identified at depth such that no flow occurs across them, a solution may be possible. However, the standard errors of estimate for hydraulic conductivity usually will be very large. Large standard errors can be common especially if measured water-level data do not accurately describe vertical-head gradients throughout the study area.

Cooley (1977) found that models of ground-water flow characteristically have large standard errors in estimates of the parameters. These errors are caused by anomalous measured water levels as well as errors in specified boundary conditions. If errors that usually occur in boundary-flux estimates are not considered during model applications, the resulting standard errors for the parameters usually will be artificially small. Test problems using the three-dimensional regression procedure supported these conclusions.

The regression model in this respect can be used to assess reliability of computed parameters and predicted values of head if the model is approximately linear with respect to the parameters. Cooley (1979) provides a test for linearity that also is applicable to three dimensions. Because the model is nonlinear with respect to hydraulic conductivity, the large number of hydraulic-conductivity parameters in most three-dimensional systems makes it more difficult to pass this test of linearity. Therefore, the use of confidence regions and test of hypothesis may not be appropriate for many three-dimensional problems. Additional testing of the regression method in three dimensions is needed to evaluate this conclusion.

#### Solute Transport

An analytical solution for one-dimensional solute transport from a specified concentration boundary into a fractured aquifer of uniform thickness is given by Bibby (1981, p. 1078). Fractures are assumed to be horizontal and separated by sufficient impermeable strata to be considered of infinite extent. The solute is considered to be nonreactive. The analytical solution was obtained by analogy to a corresponding solution for conduction and convection of heat into an aquifer.



Model results were compared to this analytical solution using representative values for aquifer characteristics. Seepage velocity of water in the fissures was set equal to 0.5 ft/d. Longitudinal and transverse dispersivities were set to 100 ft/d, and the diffusion coefficient of solute in water in the unfractured formation blocks was  $5 \times 10^{-5}$  ft<sup>2</sup>/d. The block porosity was set to unity to permit comparison with the analytical solution. The dimensions of the blocks were made sufficiently large so as not to violate assumptions in the analytical solution. Several simulations with various block dimensions were made to ensure the block dimensions were large enough not to affect the solution.

Model and analytical results after 100 days of solute transport are presented in figure 3. The close comparison verifies the model theory and program logic. The model slightly overestimates solute concentrations in the leading part of the profile and slightly underestimates concentrations in the trailing part. This characteristic probably is due to the effects of numerical dispersion associated with nodal spacing.

Additional simulations were made with the dual-porosity equation to evaluate the sensitivity of the model to variations in the dual-porosity coefficients. The coefficients that had the greatest effect on the distribution of solute appeared to be shale-matrix porosity and fracture width. The matrix-diffusion coefficient and fracture density also were important in determining the distribution of solute. These observations are in agreement with those of Grisak and Pickens (1980) and show that in fractured rock, matrix diffusion can be an important mechanism for solute transport. Factors reducing the need to consider matrix diffusion are small matrix porosities and large fracture width and density.

A review of the dual-porosity equation (eq. 5) shows that a large number of coefficients must be known even for relatively simple fracture geometries. Because in most studies these coefficients must be estimated during model calibration, the amount of water-quality data needed is far greater than the amount needed to use a continuum model. Within the framework of a three-dimensional system, sufficient data may be available only for a preliminary calibration of the model. Nevertheless in rock where matrix diffusion is dominant, a scarcity of data is not sole justification for ignoring the more complex model.

#### COMPUTER PROGRAM

A listing of the FORTRAN program that solves the three-dimensional equations of flow and solute transport in dual-porosity media is given in table 2 (in the "Supplemental Information" section at the end of the report). Data-input formats are described in table 3 (in the "Supplemental Information" section at the end of the report). Although data entry into a finite-element program typically is more cumbersome than for finite-difference programs, the increased data-entry time usually is compensated by increased flexibility in locating nodes. With a finite-element model, nodes can be accurately located at observation or pumping wells. In general, fewer nodes are needed to accurately model aquifer geometry when using a finite-element model.

Data entry into a finite-element program is more cumbersome because of the need to identify the relation among all nodes and elements. As a result, all nodes and elements must be numbered, the Cartesian coordinates of all nodes must be coded, and the nodes associated with each element must be designated.

The system used to number aquifer nodes and elements has a significant impact on the efficiency and size of the computer program. The global coefficient matrices developed in equations 2 and 4 represent the largest block of computer storage used by the program. The solution technique is more efficient, in terms of time and storage requirements, if the size of the global coefficient matrices is minimized. Storage requirements of the global coefficient matrices are directly related to the largest difference between two node numbers in an aquifer element. Therefore, efficient nodal ordering minimizes this difference and improves the efficiency of the solution. The program calculates and prints the band width. If the dimensions of arrays are not sufficient, the simulation will stop.

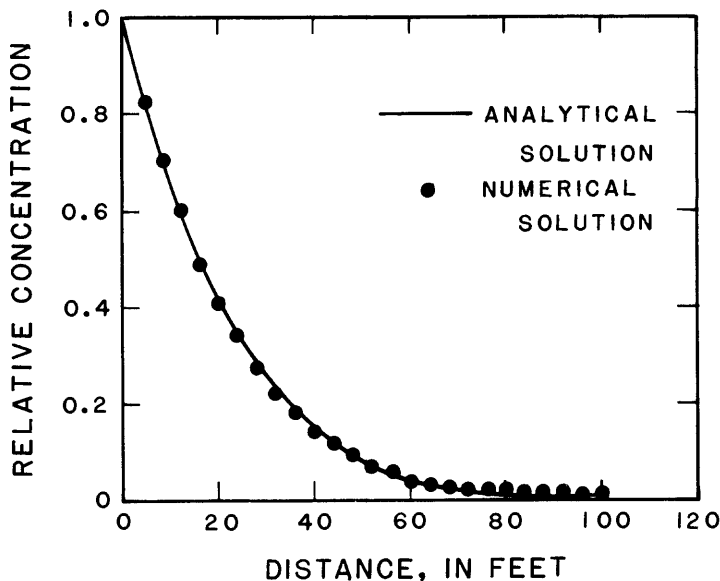


Figure 3.--Comparison of results from solute-transport model with analytical solution at 100 days.

The model procedure used to enter aquifer properties (such as hydraulic conductivity, dispersivity, or distributed recharge due to precipitation) is both flexible and easy to use. Aquifer elements are grouped into a number of user-defined aquifer zones. An aquifer property, hydraulic conductivity, for example, is simulated as the product of an element value and a zonal factor. Because hydraulic conductivity may be anisotropic, a separate zonal multiplication factor is used for each of the three principal directions. By using unique factors for each aquifer zone, it is possible to simulate varying degrees of anisotropy throughout the aquifer.

Steady-state ground-water flow properties in one or more zones can be grouped to form a single parameter of the regression analysis. Computed changes in a regression parameter are applied to the appropriate zonal factors. Element values of an aquifer property are unaffected by the regression analysis.

Boundary conditions for either ground-water flow or solute-transport may be treated as constant throughout a simulation or may be changed at user-specific times. Similarly, measured water levels and solute concentrations may be specified at any time. Input formats for transient boundary conditions and measured data are given in table 3.

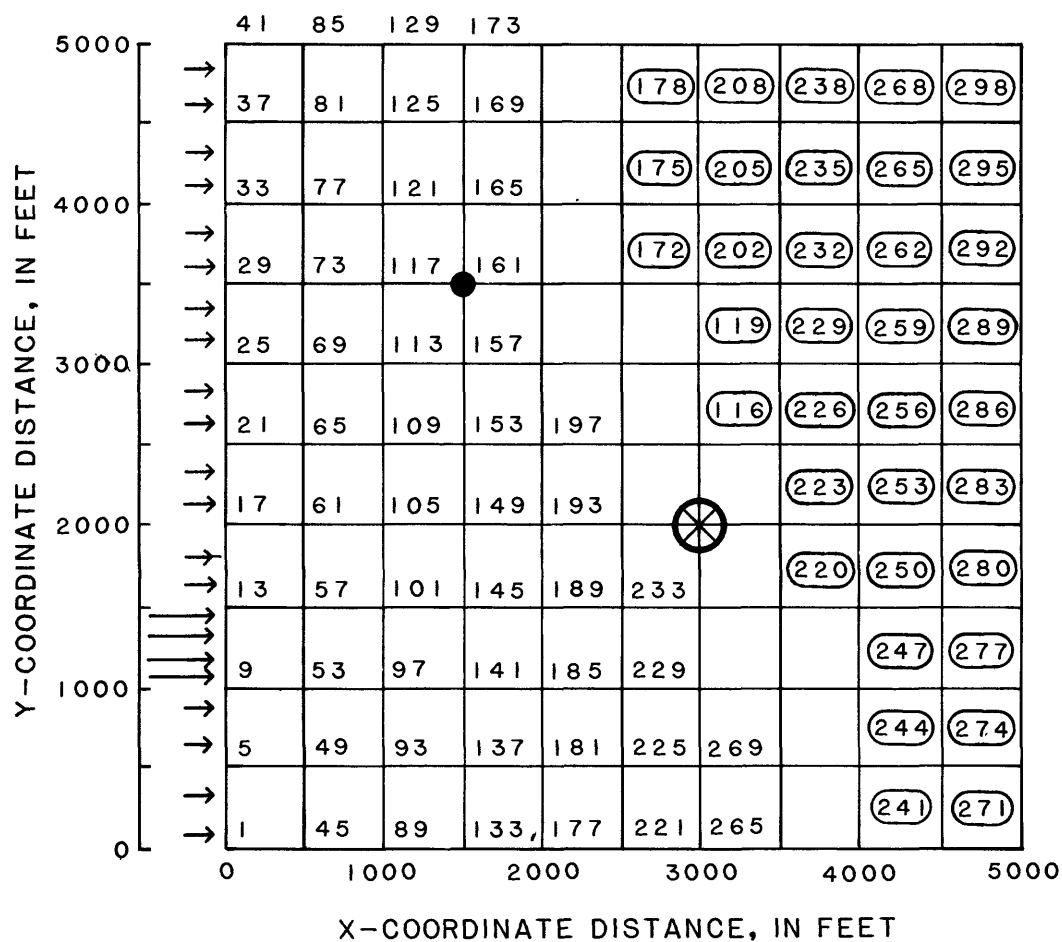
### SAMPLE SIMULATION

Use of the model is illustrated by simulating solute transport from a buried source in a three-layer aquifer system. The bottom layer represents an aquifer where horizontal joints are common within relatively thick impermeable rock. The middle layer represents an aquifer where vertical fractures are common. The top layer represents a homogeneous isotropic sandstone. Boundary conditions and nodal locations are shown horizontally in figure 4 and vertically in figure 5. Aquifer properties for each layer are given in table 1. Pumping by two wells, one from the bottom layer and one from the top, also is simulated. Steady-state flow conditions are assumed.

The input data used in the model run and results for one time step are listed in tables 4 and 5 (in the "Supplemental Information" section at the end of the report). Results present total time into the simulation, and the hydraulic head and solute concentration for each node in the finite-element grid. Principal components of Darcian velocity and dispersion coefficients also are printed for each element.

### SUMMARY

The model described in this report can simulate three-dimensional ground-water flow and solute transport in oil shale and associated hydrogeologic units. The model treats oil shale as a dual-porosity medium by simulating flow and transport within fractures using conventional finite-element methods. Diffusion of solute between fractures and the essentially static water of the shale matrix is simulated by including an analytical solution that acts as a source-sink term to the differential equation of solute transport. While knowledge of fracture orientation and spacing is needed to effectively use the model, it is not necessary to map the locations of individual fractures.



#### EXPLANATION

- |       |   |   |                          |
|-------|---|---|--------------------------|
| 45    | NODE NUMBER                                 | → | DISTRIBUTED FLUX         |
| (241) | ELEMENT NUMBER                              | → | BOUNDARY                 |
| ●     | LOCATION OF PUMPING WELL<br>IN BOTTOM LAYER | → | DISTRIBUTED FLUX AND     |
| ⊗     | LOCATION OF PUMPING WELL<br>IN TOP LAYER    | → | SOLUTE-<br>CONCENTRATION |
|       |   |   | BOUNDARY                 |

**Figure 4.--Boundary conditions and node locations along the bottom of the aquifer system used in the sample simulation.**

The computer program listed in the report incorporates many of the features of previous dual-porosity models while retaining a practical approach to solving field problems. As a result the report does not extend the theory of solute transport in any appreciable way. The emphasis in the report is on bringing together various aspects of solute-transport theory in a manner that is particularly suited to the unusual ground-water flow and solute-transport characteristics of oil-shale systems.

Methods for quantifying the uncertainty in parameter estimates that occur during model development are given in this report. The quasilinear regression method described by Cooley (1977) for estimating parameters and assessing reliability for two-dimensional models of steady-state ground-water flow has been extended to three dimensions. Because the number of model parameters in three-dimensional simulations generally is larger than in two-dimensional simulations, adequate water-level data may not be available for evaluating parameter reliability. When it is possible to evaluate parameter reliability, standard errors for hydraulic-conductivity estimates of buried strata generally are large.

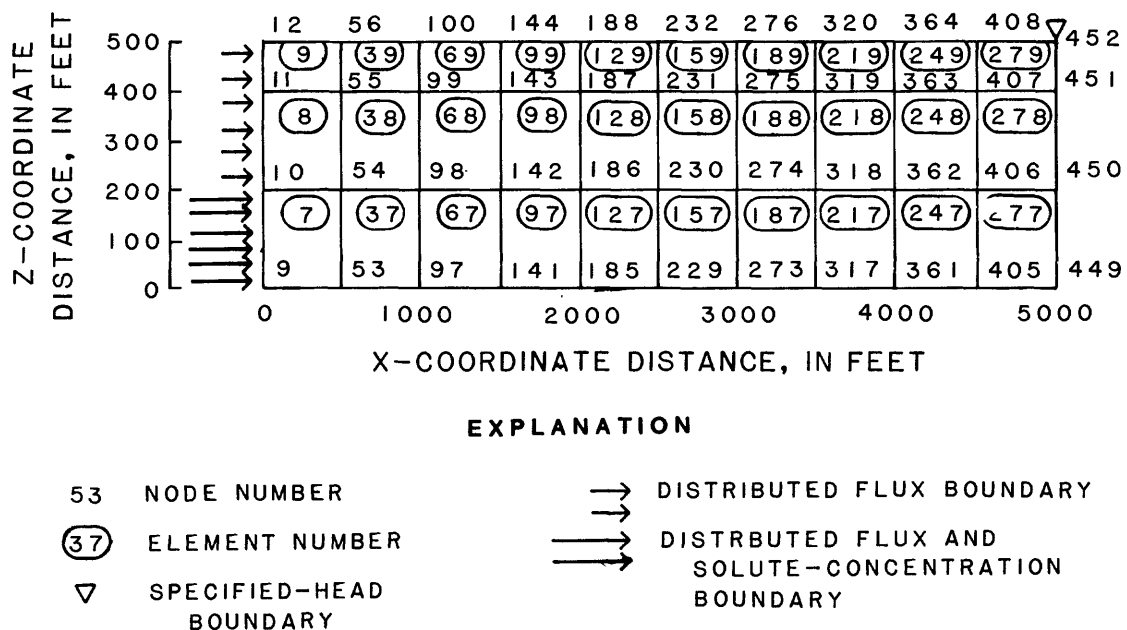


Figure 5.--Boundary conditions and node locations along a vertical section of the aquifer system used in the sample simulation.

Table 1.--Aquifer properties used in sample simulation

Aquifer property	Bottom layer	Middle layer	Top layer
Hydraulic conductivity			
x-direction (feet/day)	20.0	10.0	40.0
Hydraulic conductivity			
y-direction (feet/day)	20.0	2.0	40.0
Hydraulic conductivity			
z-direction (feet/day)	5.0	10.0	40.0
Longitudinal dispersivity (feet)	10.0	1.0	100.0
Transverse dispersivity (feet)	3.0	.3	30.0
Molecular diffusion in low permeability blocks (feet squared per day)	.0001	.0001	.0
Fracture width (feet)	.2	.01	.0
Block width (feet)	1.8	.19	.0
Effective porosity (dimensionless)	.1	.05	.3
Block porosity (dimensionless)	.01	.01	.0
Flux boundary (feet/day)	.4	.2	.4
Specified-head boundary (feet)	--	--	500.0
Solute-concentration boundary (milligrams per liter)	100.0	--	--

A large number of aquifer properties must be evaluated when simulating solute transport through dual-porosity media. Because properties such as the matrix-diffusion coefficient, matrix porosity, and fracture width and density often are estimated during model calibration, a large amount of water-quality data is needed. Within the framework of a three-dimensional system, sufficient data may be available only for a preliminary calibration of the model. Nevertheless, matrix diffusion can be an important mechanism for solute transport in fractured rock.

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**SUPPLEMENTAL INFORMATION**



Table 2.--Computer-program listing

C** A FINITE ELEMENT PROGRAM FOR THE SOLUTION OF	10
C** THREE-DIMENSIONAL TRANSIENT GROUND-WATER FLOW AND SOLUTE	20
C** TRANSPORT IN FRACTURED MATERIAL	30
C** OBSERVATIONS MAY BE VALUES OF HEAD AND CONCENTRATION IN WELLS	40
C** OPEN TO SOME OR ALL OF THE AQUIFER SYSTEM	50
C** ELEMENTS ARE ISOPARAMETRIC CUBES WITH LINEAR SIDES.	60
C** UPSTREAM FINITE ELEMENTS USED FOR SOLUTE TRANSPORT.	70
C** GUIDELINES FOR ARRAY DIMENSIONS -- IF NOT SOLVING TRANSPORT PROBLEM	80
C** THE DIMENSIONS OF TRANSPORT ARRAYS MAY BE REDUCED TO 1	90
C**     LET NVARH=NPARG+NQPAR+NBPARG	100
C**     SET IDIM.GE.NUMNP, JDIMH.GE.IBH, JDIMC.GE.IBC, KDIM.GE.NUMEL,	110
C**     NVEH.GE.NVARH	120
C** CURRENTLY DIMENSIONS ARE FOR A MAXIMUM OF	130
C** NUMNP= 490 NUMEL=300 NTIME=52 IBH= 90 IBC=180 NTPER=52 NTO=1	140
C** NTRE= 0 NTW= 0 NTS=1 NOBMAX=30 NBPARG= 0 NTB=0	150
C** NPARG= 8 NQPAR= 0 NUMAT=7 NTFRAC=9 ITRAC= 10 NLAYER= 3	160
C** MAIN ARRAYS --	170
C** XTXH(NVARH,NVARH),BH(4*NVARH)	180
C** PH(NVARH),RKH(NVARH),HINT(NTO*NOBMAX)	190
C** ZSPACE(NLAYER-1),CINT(NTO*NOBMAX),TITLE(20),HP1(NUMNP*2),	200
C** CP1(NUMNP*2),CO(NTO*NOBMAX),	210
C** NXTO(NTO+1),NXTW(NTW+1),NXTS(NTS+1),NXTB(NTB+1)	220
DIMENSION XTXH(8,8),BH(32),	230
\$         PH(8),RKH(8),HINT(30),CINT(30),	240
\$         ZSPACE(3),TITL(20),	250
\$         CO(30),NXTO(1),	260
\$         NXTW(1),NXTS(2),NXTB(1)	270
C** COFBLK ARRAYS --	280
C** H(NUMNP*2),C(NUMNP*2),QHTMP(NUMNP),WELL(NUMNP*(NTW+1)),	290
C** CFRAC(NTFRAC+1,NUMNP),CINIT(NUMNP),	300
C** WELLC(NUMNP*(NTW+1)),SFOBSH(NUMNP+NOBMAX),SFOBSC(NUMNP+NOBMAX),	310
C** WTNOD(NUMNP),DM(NUMAT),FP(NUMAT),FQ(NUMAT),	320
C** RPBR(NUMAT),CBINIT(NUMAT),RX(NUMAT),RY(NUMAT),RZ(NUMAT),	330
C** HO(NTO*NOBMAX),RC(NUMAT),RPR(NUMAT),DL(NUMAT),DT(NUMAT),	340
C** IPRMH(3,NUMAT),MAT(NUMEL),NBP(NUMNP),RMT(NUMAT)	350
COMMON/COFBLK/CFRAC(10, 490),QHTMP( 490),CINIT( 490),	360
\$         H( 980),C( 980),WELL( 490),WELLC( 490),	370
\$         RX(7),RY(7),RZ(7),RC(7),	380
\$         RPR(7),DM(7),FP(7),FQ(7),RMT(7),HO(30),	390
\$         RPBR(7),CBINIT(7),SFOBSH( 520),SFOBSC( 520),WTNOD( 490),	400
\$         DL(7),DT(7),THETAH,DTIME,TIMFAC,TIMBGN,CEQUI,DFW,	410
\$         TIMSUM,IPRMH(3,7),MAT(300),NBP( 490),IELMR	420
C** SINBLK ARRAYS --	430
C** XORD(NUMNP),YORD(NUMNP),ZORD(NUMNP),QS(NUMNP*(NTS+1)),CS(NUMNP*	440
C** (NTS+1)),QF(NQPAR),CF(NQPAR),QEH(8),QEC(8),NP(NUMEL,8),IBPRM(NUMEL)	450
COMMON/SINBLK/QEH(8),QEC(8),XORD( 490),YORD( 490),ZORD( 490),	460
\$         QS( 980),CS( 980),QF(1),CF(1),NP(300,8),IBPRM(300)	470
DOUBLE PRECISION QEH,QEC	480
C** SENBLK ARRAY --	490
C** WH(NTO*NOBMAX),WC(NTO*NOBMAX),NPBC(NUMNP,NTB+1),NODOBS(NUMNP)	500
COMMON/SENBK/WH(30),	510
\$         WC(30),THETAC,NPBC(490,1),NODOBS(490)	520

Table 2.--Computer-program listing--Continued

C** LDUBLK --	530
C** CPKH(IDIM,IBH),SH(NVARH,NUMNP),QH(NUMNP)	540
C** CPKC(IDIM,IBC),QC(IBC)	550
COMMON/LDUBLK/CPKC(490,180),CPKH(490, 90),SH(8, 490),	560
\$ QC( 490),QH( 490),IBC,IBH	570
REAL*8 CPKC,CPKH,SH,QC,QH	580
C** SHABLK,SURBLK AND LINBLK -- ALL DIMENSIONS CONSTANT	590
COMMON/SHABLK/SF(4,8,8),WF(4,8),WT(8),NUMQPT	600
COMMON/SURBLK/AF(3,4),AWT,NQPTA	610
COMMON/LINBLK/RLF(2,2),RLWT,LINOD(4,2)	620
DOUBLE PRECISION SF,WT,WF,AF,AWT	630
C** SET # RECORDS IN FILE 8 TO (NTIME/NTPER+1), # WORDS TO NUMNP	640
C** SET # RECORDS IN FILE 9 TO (NTIME+1), # WORDS TO NUMNP	650
OPEN (UNIT=8,FILE='HEADS',ACCESS='DIRECT',FORM='UNFORMATTED',	660
\$ RECL=816)	670
OPEN (UNIT=9,FILE='CONCN',ACCESS='DIRECT',FORM='UNFORMATTED',	680
\$ RECL=816)	690
OPEN (UNIT=5,FILE='SAMPLE.INPUT',STATUS='OLD')	700
OPEN (UNIT=6,FILE='SOLUTION.PRT')	710
IDIM=490	720
JDIMH= 90	730
KDIM=300	740
NVEH=8	750
LDIMH=NVEH*4	760
JDIMC=180	770
DO 105 I=1,3	780
READ(5,1)(TITL(J),J=1,20)	790
WRITE (6,2) (TITL(J),J=1,20)	800
105 CONTINUE	810
WRITE(6,3)	820
READ (5,4) NUMEL,NUMNP,NUMAT,NTIME,NTPER,NTB,NTD,NOBMAX,NTW,	830
\$ NTS,NPARH,NQPAR,NBPAR	840
WRITE(6,5) NUMEL,NUMNP,NUMAT,NTIME,NTPER,NTB,NTD,NOBMAX,NTW,	850
\$ NTS,NPARH,NQPAR,NBPAR	860
READ (5,4) NPLAYR,NEPLAY	870
WRITE (6,68) NPLAYR,NEPLAY	880
READ (5,4) ITMAX,IVELPR,IFRAC,NTFRAC,ITFRAC,NLAYER,IBEALE	890
WRITE (6,6) ITMAX,IVELPR,IFRAC,NTFRAC,ITFRAC,NLAYER,IBEALE	900
WRITE(6,7)	910
READ (5,8) NTPRT,IPO,IPRX,AP,AMP,RP,RPF,EVH	920
WRITE (6,8) NTPRT,IPO,IPRX,AP,AMP,RP,RPF,EVH	930
READ (5,9) DTIME,TIMFAC,THETAH,THETAC,DFW,CEQUI	940
WRITE (6,10) DTIME,TIMFAC,THETAH,THETAC,DFW,CEQUI	950
READ (5,9) ALFLG,ALFAX,ALFAY,ALFAZ	960
WRITE(6,11) ALFLG,ALFAX,ALFAY,ALFAZ	970
C** INITIALIZE	980
TIMBGN=DTIME	990
DO 110 I=1,NUMNP	1000
H(I)=0.0	1010
NBP(I)=0	1020
NPBC(I,1)=0	1030
NODOBS(I)=0	1040

Table 2.--Computer-program listing--Continued

110	CONTINUE	1050
	IF (THETAC.LT.0.0) GO TO 130	1060
	DO 120 I=1,NUMNP	1070
	C(I)=0.0	1080
	DO 120 J=1,NTFRAC	1090
	CFRAC(J,I)=0.0	1100
120	CONTINUE	1110
130	N=NOBMAX*NTO	1120
	NXTO(1)=999999	1130
	IF (N.LE.0) GO TO 150	1140
	DO 140 I=1,N	1150
	HO(I)=0.0	1160
	WH(I)=0.0	1170
	IF (THETAC.LT.0.0) GO TO 140	1180
	CO(I)=0.0	1190
	WC(I)=0.0	1200
140	CONTINUE	1210
150	N=NUMNP*(NTW+1)	1220
	NXTW(1)=999999	1230
	DO 156 I=1,N	1240
	WELL(I)=0.0	1250
	IF (THETAC.GE.0.0) WELLC(I)=0.0	1260
156	CONTINUE	1270
	N=NUMNP*(NTS+1)	1280
	NXTS(1)=999999	1290
	DO 158 I=1,N	1300
	QS(I)=0.0	1310
	IF (THETAC.GE.0.0) CS(I)=0.0	1320
158	CONTINUE	1330
	DO 160 I=1,NUMAT	1340
	IPRMH(1,I)=0	1350
	IPRMH(2,I)=0	1360
	IPRMH(3,I)=0	1370
160	CONTINUE	1380
	NUMM=NUMNP-NLAYER+1	1390
	NLAYM2=NLAYER-2	1400
	NLAYM1=NLAYER-1	1410
	READ (5,9) (ZSPACE(I),I=1,NLAYM1)	1420
	WRITE(6,9) (ZSPACE(I),I=1,NLAYM1)	1430
	WRITE (6,69)	1440
	DO 162 I1=1,NPLAYR	1450
	READ (5,12) I,N,XORD(I),YORD(I),	1460
	\$                  ZORD(I)	1470
	NM1=N-1	1480
	DO 162 J=1,NM1	1490
	XORD(I+J)=XORD(I)	1500
	YORD(I+J)=YORD(I)	1510
	ZORD(I+J)=ZORD(I)+ZSPACE(J)	1520
162	CONTINUE	1530
	READ (5,13) N	1540
	READ (5,14) (I,NPBC(I,1),NODOBS(I),WTNOD(I),H(I),WELL(I),QS(I),	1550
	\$                  I1=1,N)	1560

Table 2.--Computer-program listing--Continued

	WRITE(6,15)(I,NPBC(I,1),NODOBS(I),WTNOD(I),XORD(I),YORD(I),	1570
	\$ ZORD(I),H(I),WELL(I),QS(I),I=1,NUMNP)	1580
	IF (THETAC.LT.0.0) GO TO 164	1590
	WRITE (6,16)	1600
	READ (5,17) N	1610
	READ (5,17)(I,C(I),WELLC(I),CS(I),I1=1,N)	1620
	WRITE(6,67)(I,C(I),WELLC(I),CS(I),I=1,NUMNP)	1630
	DO 163 I=1,NUMNP	1640
	CINIT(I)=C(I)	1650
163	CONTINUE	1660
164	IBH=0	1670
	WRITE(6,18)	1680
	DO 170 I1=1,NEPLAY	1690
	READ (5,19) I,N,(NP(I,J),J=1,4)	1700
	NM1=N-1	1710
	MAT(I)=1	1720
	DO 166 J=5,8	1730
	NP(I,J)=NP(I,J-4)+1	1740
166	CONTINUE	1750
	IF (N1AYER.LE.2) GO TO 170	1760
	DO 168 K=1,NM1	1770
	DO 168 J=1,8	1780
	MAT(I+K)=K+1	1790
	NP(I+K,J)=NP(I,J)+K	1800
168	CONTINUE	1810
170	CONTINUE	1820
	DO 174 I=1,NUMEL	1830
	WRITE (6,20)I,(NP(I,J),J=1,8)	1840
	DO 172 J=1,8	1850
	DO 172 K=J,8	1860
	J1=IABS(NP(I,J)-NP(I,K))	1870
	IF (J1.GT.IBH) IBH=J1	1880
172	CONTINUE	1890
174	CONTINUE	1900
	IBH=IBH+1	1910
	IBC=IBH*2-1	1920
	WRITE (6,72) IBH,IBC	1930
	IF (NUMNP.GT.IDIM) STOP	1940
	IF (IBH.GT.JDIMH) STOP	1950
	IF (IBC.GT.JDIMC.AND.THETAC.GE.0.0) STOP	1960
	WRITE (6,23)	1970
	DO 188 I=1,NUMEL	1980
	IBPRM(I)=0	1990
188	CONTINUE	2000
	READ (5,4) N	2010
	IF (N.GT.0) READ (5,24)(I,MAT(I),IBPRM(I),I1=1,N)	2020
	WRITE (6,13)(I,MAT(I),IBPRM(I),I=1,NUMEL)	2030
	WRITE (6,25)	2040
	READ (5,26) (I,RX(I),RY(I),RZ(I),RC(I),I1=1,NUMAT)	2050
	WRITE (6,26) (I,RX(I),RY(I),RZ(I),RC(I),I=1,NUMAT)	2060
	ISTEDY=1	2070
	DO 190 I=1,NUMAT	2080

Table 2.--Computer-program listing--Continued

	IF (RC(I).NE.0.0) ISTEDY=0	2090
190	CONTINUE	2100
	IF (THETAC.LT.0.0) GO TO 196	2110
	WRITE (6,28)	2120
	READ (5,27) (DL(I),DT(I),DM(I),RPR(I),FP(I),FQ(I),RMT(I),	2130
	\$ RPBR(I),I=1,NUMAT)	2140
	WRITE (6,29) (I,DL(I),DT(I),DM(I),RPR(I),FP(I),FQ(I),RMT(I)	2150
	\$ ,RPBR(I),I=1,NUMAT)	2160
	IF (IFRAC.NE.1) GO TO 196	2170
	WRITE (6,32)	2180
	READ (5,27) (CBINIT(I),I=1,NUMAT)	2190
	WRITE (6,33) (CBINIT(I),I=1,NUMAT)	2200
196	IF (NPARH.LT.1) GO TO 206	2210
	DO 198 I=1,NPARH	2220
	RKH(I)=0.0	2230
198	CONTINUE	2240
	WRITE (6,34)	2250
	READ (5,35) (I,IPRMH(1,I),IPRMH(2,I),IPRMH(3,I),I1=1,NUMAT)	2260
	WRITE (6,35) (I,IPRMH(1,I),IPRMH(2,I),IPRMH(3,I),I=1,NUMAT)	2270
	IXTRFL=0	2280
	DO 200 I=1,NUMAT	2290
	IF (IPRMH(1,I).GT.0) IXTRFL=1	2300
	IF (IPRMH(2,I).GT.0) IXTRFL=1	2310
	IF (IPRMH(3,I).GT.0) IXTRFL=1	2320
200	CONTINUE	2330
	WRITE (6,36)	2340
	READ (5,37) (I,RKH(I),I1=1,NPARH)	2350
	WRITE (6,37) (I,RKH(I),I=1,NPARH)	2360
206	IF (NQPAR.LE.0) GO TO 208	2370
	WRITE (6,40)	2380
	READ (5,41) (I,QF(I),RKH(I+NPARH),I1=1,NQPAR)	2390
	WRITE (6,41) (I,QF(I),RKH(I+NPARH),I=1,NQPAR)	2400
	IXTRFL=1	2410
	IF (THETAC.LT.0.0) GO TO 208	2420
	WRITE (6,42)	2430
	READ (5,37) (I,CF(I),I1=1,NQPAR)	2440
	WRITE (6,37) (I,CF(I),I=1,NQPAR)	2450
208	IF (NBP.LE.0) GO TO 212	2460
	READ (5,43) (NBP(I),I=1,NUMNP)	2470
	WRITE(6,44) (NBP(I),I=1,NUMNP)	2480
	WRITE (6,45)	2490
	DO 210 I1=1,NBP	2500
	READ (5,37) I,RKH(I+NPARH+NQPAR)	2510
	WRITE(6,37) I,RKH(I+NPARH+NQPAR)	2520
210	CONTINUE	2530
212	NTF1=NTFRAC+1	2540
	DO 216 I=1,NUMNP	2550
	H(I+NUMNP)=H(I)	2560
	IF (THETAC.LT.0.0) GO TO 216	2570
	C(I+NUMNP)=C(I)	2580
	IF (IFRAC.EQ.0) GO TO 216	2590
	DO 214 J=1,NTF1	2600

Table 2.--Computer-program listing--Continued

214	CFRAC(J,I)=C(I)	2610
216	CONTINUE	2620
C**	READ TRANSIENT DATA	2630
	IF (NTB.LE.0.AND.ITMAX.EQ.0) GO TO 228	2640
	NXTB(1)=999999	2650
	IBCNT=1	2660
	IHR=1	2670
	ICR=1	2680
	WRITE (8'IHR) (H(J),J=1,NUMNP)	2690
	WRITE (9'ICR) (C(J),J=1,NUMNP)	2700
	IHR=IHR+1	2710
	ICR=ICR+1	2720
	NTH=NTIME/NTPER	2730
	IF (NTB.GT.0) READ(5,46) I1,I2	2740
	DO 226 I=1,NTH	2750
	IF (NTB.LE.0) GO TO 222	2760
	IF ((I-1)*NTPER.LT.I1) GO TO 222	2770
	IBM1=IBCNT	2780
	IBCNT=IBCNT+1	2790
	NXTB(IBM1)=I1	2800
	DO 218 J=1,NUMNP	2810
	NPBC(J,IBCNT)=NPBC(J,IBM1)	2820
218	CONTINUE	2830
	WRITE (6,47) I1	2840
	DO 220 K=1,I2	2850
	READ (5,48) J,NPBC(J,IBCNT),H(J),X1	2860
	IF (THETAC.GE.0.0) C(J)=X1	2870
220	CONTINUE	2880
	I1=999999	2890
	IF (IBM1.LT.NTB) READ (5,46) I1,I2	2900
222	WRITE (8'IHR) (H(J),J=1,NUMNP)	2910
	IHR=IHR+1	2920
	IF (THETAC.LT.0.0) GO TO 226	2930
	DO 224 K=1,NTPER	2940
	WRITE (9'ICR) (C(J),J=1,NUMNP)	2950
	ICR=ICR+1	2960
224	CONTINUE	2970
226	CONTINUE	2980
	NXTB(NTB+1)=999999	2990
228	IF (NTO.LE.0) GO TO 236	3000
	N=NTO*NOBMAX	3010
	DO 230 I=1,N	3020
	HINT(I)=0.0	3030
	IF (THETAC.GE.0.0) CINT(I)=0.0	3040
230	CONTINUE	3050
	WRITE (6,49)	3060
	DO 234 I=1,NTO	3070
	READ (5, 50) I1	3080
	WRITE (6, 51) I1	3090
	NXTO(I)=I1	3100
	DO 234 J=1,NOBMAX	3110
	READ (5, 52) K,X1,X2,X3,X4	3120

Table 2.--Computer-program listing--Continued

	WRITE(6, 52) K,X1,X2,X3,X4	3130
	L=NOBMAX*(I-1)+K	3140
	HO(L)=X1	3150
	WH(L)=X2	3160
232	IF (THETAC.LT.0.0) GO TO 234	3170
	CO(L)=X3	3180
	WC(L)=X4	3190
234	CONTINUE	3200
236	NXTO(NTO+1)=999999	3210
	IF (NTW.LE.0) GO TO 242	3220
	WRITE (6,53)	3230
	DO 240 I=1,NTW	3240
	DO 238 J=1,NUMNP	3250
	J1=NUMNP*I+J	3260
	WELL(J1)=WELL(J1-NUMNP)	3270
	IF (THETAC.GE.0.0) WELLC(J1)=WELLC(J1-NUMNP)	3280
238	CONTINUE	3290
	READ (5, 50) I1,I2	3300
	WRITE (6, 54) I1	3310
	NXTW(I)=I1	3320
	ITMP=MOD(I1,NTPER)	3330
	I1=NUMNP*(I)	3340
	IF (ITMP.NE.1) GO TO 248	3350
	DO 240 J=1,I2	3360
	READ (5,52) K,X1,X2	3370
	WRITE (6,52) K,X1,X2	3380
	L=I1+K	3390
	IF (ITMP.EQ.1) WELL(L)=X1	3400
	IF (THETAC.GE.0.0) WELLC(L)=X2	3410
240	CONTINUE	3420
242	NXTW(NTW+1)=999999	3430
	IF (NTS.LE.0) GO TO 250	3440
	WRITE (6,55)	3450
	DO 246 I=1,NTS	3460
	DO 244 J=1,NUMNP	3470
	J1=NUMNP*I+J	3480
	QS(J1)=QS(J1-NUMNP)	3490
	IF (THETAC.GE.0.0) CS(J1)=CS(J1-NUMNP)	3500
244	CONTINUE	3510
	READ (5, 50) I1,I2	3520
	WRITE (6, 56) I1	3530
	NXTS(I)=I1	3540
	ITMP=MOD(I1,NTPER)	3550
	I1=NUMNP*(I)	3560
	IF (ITMP.NE.1) GO TO 248	3570
	DO 246 J=1,I2	3580
	READ (5, 52) K,X1,X2	3590
	WRITE (6, 52) K,X1,X2	3600
	L=I1+K	3610
	QS(L)=X1	3620
	IF (THETAC.GE.0.0) CS(I)=X2	3630
246	CONTINUE	3640

Table 2.--Computer-program listing--Continued

	GO TO 250	3650
248	WRITE(6,80)	3660
	STOP	3670
250	NXTS(NTS+1)=999999	3680
1	FORMAT (20A4)	3690
2	FORMAT (1X,20A4)	3700
3	FORMAT (' NUMEL NUMNP NUMAT NTIME',	3710
	\$ ' NTPER NTB NTO NOBMAX NTW',	3720
	\$ ' NTS NPARH NQPAR NBPAR')	3730
4	FORMAT (14I5)	3740
5	FORMAT (14I7)	3750
6	FORMAT ('OITMAX=',I5,' IVELPR=',I5,' IFRAC=',I5,	3760
	\$ ' NTFRAC=',I5,' ITFRAC=',I5,' NLAYR=',I5,' IBEALE=',I5)	3770
7	FORMAT (' NTPRT IPO IPRX AP AMP RP RPF',	3780
	\$ ' EVH')	3790
8	FORMAT (3I5,5E10.3)	3800
9	FORMAT (7E10.3)	3810
10	FORMAT ('ODTIME = ',E12.5,' TIMFAC = ',F5.2,' THETAH = ',F5.2,	3820
	& ' THETAC = ',F5.2/' DFW = ',E12.5,	3830
	\$ ' CEQUI = ',E12.5)	3840
11	FORMAT (' ALFLG = ',E12.5,' ALFAX = ',E12.5,' ALFAY = ',E12.5,	3850
	\$ ' ALFAZ = ',E12.5)	3860
12	FORMAT (2I10,3E10.3)	3870
13	FORMAT (12I10)	3880
14	FORMAT (3I10,4E10.3)	3890
15	FORMAT (3I5,7E10.3)	3900
16	FORMAT (' NODE C WELLC CS',	3910
	\$ ' NODE C WELLC CS',	3920
	\$ ' NODE C WELLC CS')	3930
17	FORMAT (I10,3E10.3)	3940
18	FORMAT (' ELEMENT NP ARRAY')	3950
19	FORMAT (8I5)	3960
20	FORMAT (I5,5X,8I5)	3970
23	FORMAT (' ELEMENT MAT IBPRM',	3980
	\$ ' ELEMENT MAT IBPRM',	3990
	\$ ' ELEMENT MAT IBPRM',	4000
	\$ ' ELEMENT MAT IBPRM')	4010
24	FORMAT (3I10)	4020
25	FORMAT (' ZONE RX RY RZ RC')	4030
26	FORMAT (I10,4E10.3)	4040
27	FORMAT (8E10.3)	4050
28	FORMAT (' ZONE DL DT DM RP',	4060
	\$ ' FP FQ RMT RPBR')	4070
29	FORMAT (I10,8E10.3)	4080
32	FORMAT ('OCBINIT VALUES')	4090
33	FORMAT (1X,8F10.3)	4100
34	FORMAT (' ZONE IPRMH ARRAY')	4110
35	FORMAT (5I5)	4120
36	FORMAT (' PARAMETER RKH')	4130
37	FORMAT (I10,E10.3)	4140
40	FORMAT (' PARAMETER QF RKH')	4150
41	FORMAT (I10,2E10.3)	4160



Table 2.--Computer-program listing--Continued

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42  FORMAT (' PARAMETER          CF') 4170
43  FORMAT (15I5) 4180
44  FORMAT (' NBP ARRAY'/15I5) 4190
45  FORMAT (' BOUNDARY PARM      RKH') 4200
46  FORMAT (2I5) 4210
47  FORMAT ('OTIME VARYING BOUNDARY CONDITIONS -- KT = ',I5/ 4220
$    ' NODE NPBC',9X,'H',9X,'C') 4230
48  FORMAT (2I5,F10.3,F10.3) 4240
49  FORMAT ('OBSERVED HEAD AND CONCENTRATION DATA') 4250
50  FORMAT (2I10) 4260
51  FORMAT (' TIME STEP ',I5/' OBS HO WH', 4270
&    ' CO WC') 4280
52  FORMAT (I10,4E10.3) 4290
53  FORMAT ('OTIME VARYING POINT SOURCE-SINK DATA') 4300
54  FORMAT (' TIME STEP ',I5/' NODE WELL WELLC') 4310
55  FORMAT ('OTIME VARYING LINE SOURCE-SINK DATA') 4320
56  FORMAT (' TIME STEP ',I5/' NODE QS CS') 4330
57  FORMAT (' TIME STEP ',I5,' TOTAL TIME OF SIMULATION',F12.2/ 4340
$    ' NODE XORD YORD', 4350
$    ' ZORD H C') 4360
58  FORMAT (I10,5E12.5) 4370
59  FORMAT (I10,3E12.5,12X,E12.5) 4380
60  FORMAT ('O OBS HO HINT', 4390
$    ' CO CINT') 4400
61  FORMAT (I5,4(1X,E12.5)) 4410
62  FORMAT (I5,26X,2(1X,E12.5)) 4420
63  FORMAT ('OSTATISTICS OF INITIAL HEAD SOLUTION') 4430
64  FORMAT ('OSTATISTICS OF INITIAL CONCENTRATION SOLUTION') 4440
65  FORMAT ('ONUMBER OF OBSERVATIONS = ',I5/' ESTIMATED SUM OF ', 4450
$    'SQUARED ERRORS FOR INITIAL SOLUTION = ',E12.5/ 4460
$    ' ERROR VARIANCE FOR INITIAL SOLUTION = ',E12.5) 4470
66  FORMAT ('OINITIAL SOLUTION') 4480
67  FORMAT (I10,3E10.3,I10,3E10.3,I10,3E10.3) 4490
68  FORMAT (' NPLAYR=',I5,' NEPLAY=',I5) 4500
69  FORMAT (' NODE NPBC NODOBS WINOD XORD YORD ZORD', 4510
$    ' H WELL QS') 4520
70  FORMAT (' FLOW PROBLEM') 4530
71  FORMAT (' SOLUTE TRANSPORT PROBLEM') 4540
72  FORMAT ('OFLOW PROBLEM BAND WIDTH = ',I5/' TRANSPORT PROBLEM ', 4550
$    'BAND WIDTH = ',I5) 4560
73  FORMAT (' UPDATED PARAMETERS'/' MATI RX', 4570
$    ' RY RZ ') 4580
74  FORMAT (I5,3(1X,E11.5)) 4590
75  FORMAT (' PARAMETER QF ') 4600
76  FORMAT (I10,1X,E11.5) 4610
77  FORMAT ('FLOW PARAMETER ',I5,' EFFECTIVELY ZERO') 4620
78  FORMAT (' SOLUTION FAILED TO CONVERGE IN ',I5,' ITERATIONS') 4630
79  FORMAT (' SOLUTION CONVERGED IN ',I5,' ITERATIONS') 4640
80  FORMAT (' TIME STEP IS INVALID FOR A CHANGE OF PARAMETERS'/ 4650
$    ' TIME STEP MUST BE DIVISIBLE BY NTPER WITH A REMAINDER', 4660
$    ' OF 1') 4670
CALL SHAFAC(THETAC) 4680

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Table 2.--Computer-program listing--Continued

CALL SURFAC	4690
CALL LINFAC	4700
NVARH=NPARH+NQPAR+NBPAP	4710
NVHX2=NVARH+NVARH	4720
NVHX3=NVARH+NVHX2	4730
NODES=8	4740
C** COMPUTE AND COUNT PRIOR INFORMATION	4750
NPRIRH=0	4760
IF (NVARH.LT.1) GO TO 254	4770
DO 252 I=1,NVARH	4780
PH(I)=1.0	4790
IF (RKH(I).LE.0.0) GO TO 252	4800
RKH(I)=EVH/(RKH(I)*RKH(I))	4810
NPRIRH=NPRIRH+1	4820
252 CONTINUE	4830
C** COMPUTE INITIAL SOLUTION	4840
254 WRITE (6,66)	4850
IFLOW=0	4860
KTH=0	4870
IOBCNT=1	4880
IWCNT=1	4890
ISCNT=1	4900
IBCNT=1	4910
SKBIG=0.0	4920
SKBIG2=0.0	4930
DTIME=TIMBGN/TIMFAC	4940
TIMSUM=0.0	4950
OBSh=0.0	4960
YSQH=0.0	4970
ERVARH=0.0	4980
OBSC=0.0	4990
YSQC=0.0	5000
ERVARC=0.0	5010
IOBCNT=1	5020
DO 294 KT=1,NTIME	5030
DTIME=DTIME*TIMFAC	5040
TIMSUM=TIMSUM+DTIME	5050
LU=0	5060
IFLOW=IFLOW+1	5070
IF (IFLOW.GT.NTPER) IFLOW=1	5080
IF (IFLOW.EQ.1) KTH=KTH+1	5090
IF (KT.EQ.1) LU=1	5100
IF (KT.GT.NXTO(IOBCNT)) IOBCNT=IOBCNT+1	5110
IF (KT.EQ.NXTB(IBCNT)) LU=1	5120
IF (TIMFAC.GT.1.01) LU=1	5130
LUH=LU	5140
LUC=LU	5150
IF (KT.EQ.NXTW(IWCNT)) IWCNT=IWCNT+1	5160
IF (KT.EQ.NXTS(ISCNT)) ISCNT=ISCNT+1	5170
IF (KT.EQ.NXTB(IBCNT)) IBCNT=IBCNT+1	5180
IF (IFLOW.GT.1) GO TO 260	5190
IF (NTB.LE.0) GO TO 258	5200

Table 2.--Computer-program listing--Continued

	IHR=KTH+1	5210
	DO 256 I=1,NUMNP	5220
	J=I+NUMNP	5230
	H(I)=H(J)	5240
256	CONTINUE	5250
	J1=NUMNP+1	5260
	J2=NUMNP*2	5270
	READ (8'IHR) (H(J),J=J1,J2)	5280
258	CALL FLOW (IWCNT,ISCNT,IBCNT,KT,KTH,	5290
	\$ NTPER,NUMNP,NVARH,NPARH,LUH,NUMEL,	5300
	\$ ISTEDY,0,IDIM)	5310
	CALL UDU (NUMNP,LUH)	5320
260	ICR=KT+1	5330
	DO 262 I=1,NUMNP	5340
	J=I+NUMNP	5350
	IF (IFLOW.EQ.1) H(J)=QH(I)	5360
	IF (THETAC.LT.0.0) GO TO 262	5370
	C(I)=C(J)	5380
262	CONTINUE	5390
	IF (THETAC.LT.0.0) GO TO 270	5400
	IF (NTB.LE.0) GO TO 264	5410
	J1=NUMNP+1	5420
	J2=NUMNP*2	5430
	READ(9'ICR) (C(J),J=J1,J2)	5440
264	IVPTMP=0	5450
	IF (MOD(KT,NTPRT).EQ.0) IVPTMP=1	5460
	CALL SALT (IFLOW,IWCNT,ISCNT,IBCNT,KT,KTH,	5470
	\$ NTPER,NUMNP,LUC,NUMEL,ISTEDY,IVPTMP,IFRAC,1,0,	5480
	\$ IDIM,ALFAX,ALFAY,ALFAZ,ALFLG,NTFRAC,ITFRAC,	5490
	\$ NOBMAX,IOBCNT)	5500
	CALL LDU (NUMNP,LUC)	5510
	DO 268 I=1,NUMNP	5520
	I1=NUMNP+I	5530
	C(I1)=QC(I)	5540
	IF (IFRAC.EQ.0) GO TO 268	5550
	DO 266 J=2,NTF1	5560
	JM1=J-1	5570
	CFRAC(JM1,I)=CFRAC(J,I)	5580
266	CONTINUE	5590
	CFRAC(NTF1,I)=QC(I)	5600
268	CONTINUE	5610
270	IHR=KTH+1	5620
	ICR=KT+1	5630
	J1=NUMNP+1	5640
	J2=NUMNP*2	5650
	IF (IFLOW.GT.1.OR.NVARH.EQ.0) GO TO 272	5660
	WRITE (8'IHR) (H(J),J=J1,J2)	5670
272	IF (THETAC.LT.0.0) GO TO 274	5680
	WRITE (9'ICR) (C(J),J=J1,J2)	5690
274	IF (MOD(KT,NTPRT).NE.0) GO TO 280	5700
	WRITE (6,57) KT,TIMSUM	5710
	DO 278 I=1,NUMNP	5720

Table 2.--Computer-program listing--Continued

	I1=NUMNP+I	5730
	I2=NUMNP+I	5740
	IF (THETAC.GE.0.0) GO TO 276	5750
	IF (IFLOW.EQ.1) WRITE (6,58) I,XORD(I),YORD(I),ZORD(I),H(I2)	5760
	GO TO 278	5770
276	IF (IFLOW.EQ.1) WRITE (6,58) I,XORD(I),YORD(I),ZORD(I),H(I2),C(I1)	5780
	IF (IFLOW.NE.1) WRITE (6,59) I,XORD(I),YORD(I),ZORD(I),C(I1)	5790
278	CONTINUE	5800
280	IF (NOBMAX.LE.0) GO TO 294	5810
	IF (KT.NE.NXTO(IOBCNT)) GO TO 294	5820
C**	COMPUTE INITIAL ERROR VARIANCE	5830
	DO 286 I=1,NUMNP	5840
282	IF (NODOBS(I).LE.0) GO TO 286	5850
	I3=NUMNP+I	5860
	I2=NODOBS(I)+NOBMAX*(IOBCNT-1)	5870
	IF (IFLOW.GT.1) GO TO 284	5880
	HINT(I2)=HINT(I2)+SFOBSH(I)*H(I3)	5890
284	IF (THETAC.LT.0.0) GO TO 286	5900
	CINT(I2)=CINT(I2)+SFOBSC(I)*C(I3)	5910
286	CONTINUE	5920
	WRITE (6,60)	5930
	DO 290 I=1,NOBMAX	5940
	I2=I+NOBMAX*(IOBCNT-1)	5950
	IF (IFLOW.EQ.1.AND.THETAC.GE.0.0) WRITE (6,61) I,HO(I2),HINT(I2)	5960
\$	,CO(I2),CINT(I2)	5970
	IF (IFLOW.EQ.1.AND.THETAC.LT.0.0) WRITE (6,61) I,HO(I2),HINT(I2)	5980
	IF (IFLOW.GT.1.AND.THETAC.GE.0.0) WRITE (6,62) I,CO(I2),CINT(I2)	5990
	IF (IFLOW.GT.1) GO TO 288	6000
	IF (WH(I2).LE.0.0) GO TO 288	6010
	OBSH=OBSH+1.0	6020
	YSQH=YSQH+WH(I2)*(HO(I2)-HINT(I2))**2	6030
288	IF (THETAC.LT.0.0) GO TO 290	6040
	IF (WC(I2).LE.0.0) GO TO 290	6050
	OBSC=OBSC+1.0	6060
	YSQC=YSQC+WC(I2)*(CO(I2)-CINT(I2))**2	6070
290	CONTINUE	6080
	IOBCNT=IOBCNT+1	6090
292	CONTINUE	6100
294	CONTINUE	6110
	IF (NOBMAX.LE.0) GO TO 640	6120
	ERVARH=YSQH/(OBSH-NVARH+NPRIRH)	6130
	NTMP=OBSH	6140
	WRITE (6,63)	6150
	WRITE (6,65) NTMP,YSQH,ERVARH	6160
296	IF (THETAC.LT.0.0) GO TO 298	6170
	ERVARC=YSQC/(OBSC-NVARC+NPRIRC)	6180
	NTMP=OBSC	6190
	WRITE (6,64)	6200
	WRITE (6,65) NTMP,YSQC,ERVARC	6210
C**	BEGIN ITERATIONS	6220
298	INDT=0	6230
	ER=0.01	6240

Table 2.--Computer-program listing--Continued

	ERP=1000.0	6250
	IF (ITMAX.LE.0) GO TO 640	6260
	IF (ISTEDY.EQ.0) GO TO 640	6270
	DO 372 ITER=1,ITMAX	6280
	IOBCNT=1	6290
	IWCNT=1	6300
	ISCNT=1	6310
	IBCNT=1	6320
	IFLOW=1	6330
	KTH=1	6340
	DTIME=TIMBGN/TIMFAC	6350
	TIMSUM=0.0	6360
	IF (AMP.LT.-.5) GO TO 304	6370
	YSQH=0.0	6380
	DO 302 I=1,NVARH	6390
	BH(I)=0.0	6400
	DO 302 J=1,NVARH	6410
	XTXH(I,J)=0.0	6420
302	CONTINUE	6430
304	IHR=1	6440
	J1=NUMNP+1	6450
	J2=NUMNP*2	6460
	READ (8'IHR) (H(J),J=J1,J2)	6470
	DTIME=DTIME*TIMFAC	6480
	TIMSUM=TIMSUM+DTIME	6490
	LU=0	6500
	DO 316 I=1,NUMNP	6510
	J=I+NUMNP	6520
	H(I)=H(J)	6530
316	CONTINUE	6540
	J1=NUMNP+1	6550
	J2=NUMNP*2	6560
	IHR=KTH+1	6570
	READ (8'IHR)(H(J),J=J1,J2)	6580
	CALL FLOW (IWCNT,ISCNT,IBCNT,KT,KTH,	6590
	\$ NTPER,NUMNP,NVARH,NPARH,LU,NUMEL,	6600
	\$ ISTEDY,ITER,IDIM)	6610
	CALL SENS (NVARH,NUMNP,IPO,KT,KTH,IOBCNT,	6620
	\$ IXTRFL,NOBMAX,IBCNT)	6630
	DO 322 I=1,NUMNP	6640
	QHTMP(I)=QH(I)	6650
322	CONTINUE	6660
	DO 324 I=1,NUMNP	6670
	IF (NPBC(I,IBCNT).GE.0.OR.NPBC(I,IBCNT).EQ.-2.OR.NPBC(I,IBCNT)	6680
	\$ .EQ.-4) GO TO 324	6690
	IF (NBP(I).LE.0) GO TO 324	6700
	N=NBP(I)+NPARH+NQPAR	6710
	J=NUMNP+I	6720
	SH(N,I)=H(J)*THETAH+(1.0-THETAH)*H(I)	6730
324	CONTINUE	6740
330	CONTINUE	6750
C**	FORM LEAST SQUARES MATRIX	6760

Table 2.--Computer-program listing--Continued

	IF (AMP.LT.-.5) GO TO 344	6770
	DO 342 I=1,NOBMAX	6780
	I1=I+(IOBCNT-1)*NOBMAX	6790
	HTMP=0.0	6800
	DO 332 J=1,NUMNP	6810
	IF (NODOBS(J).NE.I) GO TO 332	6820
	HTMP=HTMP+SFOBSH(J)*QHTMP(J)	6830
332	CONTINUE	6840
	IF (WH(I1).LE.0.0) GO TO 342	6850
	TEMP=HO(I1)-HTMP	6860
	DO 340 K=1,NVARH	6870
	STMP=0.0	6880
	DO 334 J=1,NUMNP	6890
	IF (NODOBS(J).EQ.I) STMP=STMP+SFOBSH(J)*SH(K,J)	6900
334	CONTINUE	6910
	TMP=WH(I1)*STMP	6920
	DO 338 L=1,NVARH	6930
	STMP=0.0	6940
	DO 336 J=1,NUMNP	6950
	IF (NODOBS(J).EQ.I) STMP=STMP+SFOBSH(J)*SH(L,J)	6960
336	CONTINUE	6970
338	XTXH(L,K)=XTXH(L,K)+TMP*STMP	6980
340	BH(K+NVHX2)=BH(K+NVHX2)+TMP*TEMP	6990
	YSQH=YSQH+TEMP*TEMP*WH(I1)	7000
342	CONTINUE	7010
344	CONTINUE	7020
	J1=NUMNP+1	7030
	J2=NUMNP*2	7040
	IHR=KTH+1	7050
	IF (IFLOW.EQ.1) WRITE (8'IHR) (H(J),J=J1,J2)	7060
	WRITE (6,70)	7070
	CALL LSTSQ (XTXH,BH,RKH,PH,YSQH,AP,AMP,RP,RPF,NVARH,NVHX2,NVHX3,	7080
	\$ IPO,INDT,ITER,NVEH,LDIMH)	7090
	IF (INDT.EQ.1) GO TO 521	7100
C**	UPDATE HEAD	7110
348	DO 352 I=1,NUMNP	7120
	I1=NUMNP+I	7130
	SUMH=0.0	7140
	IF (NVARH.LE.0) GO TO 352	7150
	DO 350 J=1,NVARH	7160
	SUMH=SUMH+BH(J)*SH(J,I)	7170
350	CONTINUE	7180
	IF (IFLOW.EQ.1) H(I1)=QHTMP(I)+SUMH	7190
352	CONTINUE	7200
	IHR=2	7210
	J1=NUMNP+1	7220
	J2=2*NUMNP	7230
	WRITE (8'IHR) (H(J),J=J1,J2)	7240
C**	UPDATE PARAMETERS	7250
354	IF (IPO.EQ.1) WRITE (6,73)	7260
	DO 356 I=1,NUMAT	7270
	IF (NVARH.LE.0) GO TO 356	7280

Table 2.--Computer-program listing--Continued

L=IPRMH(1,I)	7290
IF (L.GT.0) RX(I)=RX(I)*(BH(L)+1.0)	7300
L=IPRMH(2,I)	7310
IF (L.GT.0) RY(I)=RY(I)*(BH(L)+1.0)	7320
L=IPRMH(3,I)	7330
IF (L.GT.0) RZ(I)=RZ(I)*(BH(L)+1.0)	7340
IF (IPO.EQ.1) WRITE (6,74) I,RX(I),RY(I),RZ(I)	7350
356 CONTINUE	7360
C** UPDATE SURFACE AND POINT SOURCE-SINK PARAMETERS	7370
IF (NQPAR.LE.0) GO TO 360	7380
WRITE (6,75)	7390
DO 358 I=1,NQPAR	7400
QF(I)=QF(I)*(BH(I+NPARH)+1.0)	7410
IF (IPO.EQ.1) WRITE (6,76) I,QF(I)	7420
358 CONTINUE	7430
360 CONTINUE	7440
ITMP=NUMNP*NTIME/NTPER	7450
C** CHECK FOR CONVERGENCE	7460
362 DO 364 I=1,NVARH	7470
IF (ABS(BH(I)/AP).GT.ER) GO TO 366	7480
364 CONTINUE	7490
GO TO 374	7500
C** CALCULATE NEW SCALED PRIOR INFORMATION PARAMETERS	7510
366 IND=0	7520
DO 370 I=1,NVARH	7530
TEMP=BH(I)+1.0	7540
PH(I)=PH(I)/TEMP	7550
IF (ABS(PH(I)).LT.ERP) GO TO 368	7560
WRITE (6, 77) I	7570
IND=1	7580
368 RKH(I)=RKH(I)*TEMP*TEMP	7590
370 CONTINUE	7600
C** END ITERATION LOOP	7610
372 CONTINUE	7620
WRITE (6, 78) ITMAX	7630
GO TO 376	7640
374 WRITE (6, 79) ITER	7650
376 CONTINUE	7660
C** COMPUTE SUM OF SQUARES AND CORRELATION COEFFICIENT	7670
378 SUMA=0.	7680
SUMB=0.	7690
SUMC=0.	7700
SUMD=0.	7710
SUM=0.	7720
DO 382 N=1,NOBMAX	7730
IF (WH(N).LE.0.) GO TO 382	7740
TMP=WH(N)**.5	7750
WH(N)=TMP	7760
TEMP=TMP*HO(N)	7770
HTMP=0.0	7780
DO 380 J=1,NUMNP	7790
IF (NODOBS(J).EQ.N) HTMP=HTMP+SFOBSH(J)*H(J+NUMNP)	7800

Table 2.--Computer-program listing--Continued

380	CONTINUE	7810
	TMP=TMP*HTMP	7820
	SUMA=SUMA+TEMP	7830
	SUMB=SUMB+TMP	7840
	SUMC=SUMC+TEMP*TEMP	7850
	SUMD=SUMD+TMP*TMP	7860
	SUM=SUM+TEMP*TMP	7870
382	CONTINUE	7880
	RCOEF=(OBESH*SUM-SUMA*SUMB)/((OBESH*SUMC-SUMA*SUMA)*	7890
	\$ (OBESH*SUMD-SUMB*SUMB)**0.5	7900
	TEMP=AP*(2.-AP)	7910
	TPA=2.*AP*RP/TEMP	7920
	SUM=0.0	7930
	DO 384 I=1,NVARH	7940
	TMP=PH(I)-1.0	7950
	SUM=SUM+TEMP*BH(I+NVARH)*(BH(I+NVHX2)-TPA*BH(I+NVHX3))*	7960
	\$ (RPF*PH(I)-1.0))-RKH(I)*TMP*TMP	7970
384	CONTINUE	7980
	YSQH=YSQH-SUM	7990
C**	COMPUTE SCALED VARIANCE-COVARIANCE MATRIX, ERROR VARIANCE, AND	8000
C**	OPTIMUM BIAS PARAMETER:	8010
C**	CORRECT XTXH FOR MARQUARDT PARAMETER	8020
	IF (NVARH.EQ.1) GO TO 435	8030
	IF(AMP.LE.0) GO TO 390	8040
	DO 388 I=1,NVARH	8050
	XTXH(I,I)=1.+RP	8060
	DO 386 J=1,NVARH	8070
386	XTXH(J,I)=XTXH(I,J)	8080
388	CONTINUE	8090
	AMP=-1.0	8100
	CALL LSTSQ (XTXH,BH,RKH,PH,YSQH,AP,AMP,RP,RPF,NVARH,NVX2,NVX3,	8110
	\$ IPO,INDT,ITER,NVEH,LDIMH)	8120
	IF(INDT.GT.0) GO TO 521	8130
C**	COMPUTE XTXH-INVERSE	8140
390	XTXH(NVARH,NVARH)=1.0/XTXH(NVARH,NVARH)	8150
	NM1=NVARH-1	8160
	DO 430 K=1,NM1	8170
	KP1=K+1	8180
	DO 400 I=KP1,NVARH	8190
	SUM=0.0	8200
	IM1=I-1	8210
	DO 392 J=K,IM1	8220
392	SUM=SUM+XTXH(I,J)*XTXH(J,K)	8230
	XTXH(K,I)=-SUM	8240
400	XTXH(I,K)=-SUM*XTXH(I,I)	8250
	DO 420 J=1,K	8260
	SUM=XTXH(K,J)	8270
	DO 410 I=KP1,NVARH	8280
410	SUM=SUM+XTXH(I,J)*XTXH(K,I)	8290
	XTXH(K,J)=SUM	8300
420	XTXH(J,K)=XTXH(K,J)	8310
430	CONTINUE	8320



Table 2.--Computer-program listing--Continued

	GO TO 440	8330
435	XTXH(1,1)=1.0/(1.+RP)	8340
440	DO 450 J=1,NVARH	8350
	BH(J)=0.0	8360
	BH(J+NVARH)=AP*BH(J+NVARH)+BH(J+NVHX3)	8370
	PH(J)=BH(J+NVHX3)*PH(J)	8380
450	XTXH(J,NVARH)=XTXH(NVARH,J)	8390
C**	COMPUTE TR(XTXH-INVERSE**2) AND RPF	8400
	TRACE=0.0	8410
	RPF=0.0	8420
	IF(RP.LE.0.0) GO TO 459	8430
	DO 457 N=1,NVARH	8440
	DO 452 J=1,NVARH	8450
452	WELL(J)=XTXH(J,N)	8460
	SUMA=0.	8470
	DO 456 J=N,NVARH	8480
	SUM=0.	8490
	DO 454 I=1,NVARH	8500
454	SUM=SUM+WELL(I)*XTXH(I,J)	8510
	QHTMP(J)=SUM	8520
	SUMA=SUMA+PH(J)*SUM	8530
	BH(J)=BH(J)+PH(N)*SUM	8540
456	XTXH(J,N)=XTXH(J,N)-RP*SUM	8550
	BH(N)=BH(N)+SUMA-QHTMP(N)*PH(N)	8560
457	TRACE=TRACE+QHTMP(N)	8570
	IF(IRPF.LE.0) GO TO 459	8580
	SUM=0.	8590
	DO 458 I=1,NVARH	8600
	RPF=RPF+BH(I)*BH(I+NVARH)	8610
458	SUM=SUM+BH(I)*PH(I)	8620
	RPF=RPF/SUM	8630
C**	COMPUTE ERROR VARIANCE AND COV(SCALED PARAMETERS)	8640
459	TEMP=NVARH-NPRIRH	8650
	VARH=YSQH/(OBSh-TEMP+RP*RP*TRACE)	8660
	SUM=0.0	8670
	DO 462 J=1,NVARH	8680
	TMP=BH(J+NVARH)-RPF*PH(J)	8690
	SUM=SUM+TMP*TMP	8700
	TEMP=BH(J+NVHX3)	8710
	DO 460 I=J,NVARH	8720
	XTXH(I,J)=VARH*XTXH(I,J)/(BH(I+NVHX3)*TEMP)	8730
460	XTXH(J,I)=XTXH(I,J)	8740
462	BH(J)=XTXH(J,J)**0.5	8750
	TEMP=NVARH	8760
	SUM=TEMP*VARH/SUM	8770
C**	PRINT ERROR VARIANCE, ESTIMATED SUM OF SQUARED ERRORS, CORRELATION	8780
C**	COEFFICIENT, AND OPTIMUM RIDGE PARAMETERS	8790
	WRITE(6,864) VARH,YSQH,RCOEF,SUM,RPF	8800
C**	PRINT SCALED VARIANCE-COVARIANCE MATRIX	8810
	WRITE(6,874)	8820
	CALL PRTOT(XTXH,NVEH,0,NVARH)	8830
C**	COMPUTE AND PRINT CORRELATION MATRIX FOR PARAMETERS	8840

Table 2.--Computer-program listing--Continued

	DO 510 J=1,NVARH	8850
	TEMP=BH(J)	8860
	DO 500 I=J,NVARH	8870
	XTXH(I,J)=XTXH(I,J)/(BH(I)*TEMP)	8880
500	XTXH(J,I)=XTXH(I,J)	8890
510	CONTINUE	8900
	WRITE(6,876)	8910
	CALL PRTOT(XTXH,NVEH,0,NVARH)	8920
521	IF(IPRX.LE.0.AND.ITER.LT.ITMAX.AND.IBEALE.EQ.0) STOP	8930
	IF(NVAR.LT.2.AND.IBEALE.EQ.0) STOP	8940
C**	ORTHOGONALIZE SENSITIVITY MATRIX, COMPRESS SH	8950
	N=NOBMAX	8960
	DO 536 I=1,NVARH	8970
	DO 532 J=1,NOBMAX	8980
	WH(J)=0.0	8990
532	CONTINUE	9000
	DO 534 J=1,NUMNP	9010
	K=NODOBS(J)	9020
	IF (K.GT.0) WH(K)=WH(K)+SFOBSH(J)*SH(I,J)	9030
534	CONTINUE	9040
	DO 535 J=1,NOBMAX	9050
	SH(I,J)=WH(J)	9060
535	CONTINUE	9070
536	CONTINUE	9080
	IF(NPRIRH.LE.0) GO TO 539	9090
	DO 538 I=1,NVARH	9100
	IF(RKH(I).LT.1.E-10) GO TO 538	9110
	N=N+1	9120
	DO 537 J=1,NVARH	9130
537	SH(J,N)=0.0	9140
	SH(I,N)=RKH(I)**0.5	9150
538	CONTINUE	9160
C**	ORTHOGONALIZE SH	9170
539	IF (IBEALE.EQ.0) GO TO 542	9180
	WRITE (6,860)	9190
	DO 540 I=1,NOBMAX	9200
	WRITE (6,862) I,(SH(J,I),J=1,NVARH)	9210
540	CONTINUE	9220
	IF (IPRX.LE.0.AND.ITER.LT.ITMAX) STOP	9230
	IF (NVAR.LT.2) STOP	9240
542	NTMP=OBESH+NPRIRH	9250
	DO 544 I=1,NTMP	9260
544	WH(I)=SH(1,I)	9270
	DO 600 N=2,NVARH	9280
	NM1=N-1	9290
	SUM=0.0	9300
	DO 550 I=1,NTMP	9310
	SUM=SUM+WH(I)*WH(I)	9320
	SH(NM1,I)=WH(I)	9330
550	CONTINUE	9340
	IF(SUM.LT.1.E-20) GO TO 610	9350
	BH(NM1)=1.0/SUM	9360

Table 2.--Computer-program listing--Continued

```

DO 570 J=1,NM1                      9370
SUM=0.0                              9380
DO 560 K=1,NTMP                      9390
560 SUM=SUM+BH(J)*SH(J,K)*SH(N,K)    9400
570 WELL(J)=SUM                      9410
DO 590 K=1,NTMP                      9420
SUM=0.0                              9430
DO 580 I=1,NM1                      9440
580 SUM=SUM+SH(I,K)*WELL(I)          9450
590 WH(K)=SH(N,K)-SUM               9460
600 CONTINUE                        9470
610 WRITE(6,886)                    9480
K=1                                  9490
L=8                                  9500
DO 630 M=1,NVARH,8                  9510
IF(L.GT.NVARH) L=NVARH              9520
WRITE(6,888) (I,I=K,L)              9530
DO 620 J=1,NTMP                    9540
SH(NVARH,J)=WH(J)                   9550
WRITE(6,890) J,(SH(I,J),I=K,L)      9560
620 CONTINUE                        9570
K=K+8                               9580
L=L+8                               9590
630 CONTINUE                        9600
640 CONTINUE                        9610
STOP                                9620
860 FORMAT (' OBS',10X,'DEPTH-INTEGRATED SENSITIVITY COEFFICIENTS') 9630
862 FORMAT (15,9E14.7,1X/5X,9E14.7,1X/5X,9E14.7) 9640
864 FORMAT ('OERROR VARIANCE = ',E12.5/' ESTIMATED SUM OF SQUARED ', 9650
$'ERRORS = ',E12.5/' CORRELATION COEFFICIENT = ',F6.4/' OPTIMUM RID 9660
$GE PARAMETER = ',E12.5/' OPTIMUM BIAS PARAMETER = ',E12.5) 9670
874 FORMAT ('OSCALED VARIANCE-COVARIANCE MATRIX') 9680
876 FORMAT ('OCORRELATION MATRIX FOR PARAMETERS') 9690
882 FORMAT ('OSCALED SENSITIVITY ARRAYS') 9700
884 FORMAT('OPARAMETER NUMBER ',I5) 9710
886 FORMAT(1H0,5X,'ORTHOGONALIZED SENSITIVITY MATRIX'/ 9720
$ ' OBSERVATION NUMBER AND VALUES FOR EACH PARAMETER') 9730
888 FORMAT (15X,'PARAMETER NOS. '/16X,8(I3,9X)) 9740
890 FORMAT (2X,I4,8(1X,G11.5)) 9750
END 9760
SUBROUTINE FLOW (IWCNT,ISCNT,IBCNT,KT, 9770
$ KTH,NTPER,NUMNP,NVARH,NPARH,LU, 9780
$ NUMEL,ISTEDY,ITER,IDIM) 9790
C** ROUTINE TO FORMULATE GLOBAL STIFFNESS AND 9800
C** SENSITIVITY MATRICES 9810
COMMON/COFBLK/CFRAC(10, 490),QHTMP( 490),CINIT( 490), 9820
$ H( 980),C( 980),WELL( 490),WELLC( 490), 9830
$ RX(7),RY(7),RZ(7),RC(7), 9840
$ RPR(7),DM(7),FP(7),FQ(7),RMT(7),HO(30), 9850
$ RPBR(7),CBINIT(7),SFOBSH( 520),SFOBSC( 520),WTNOD( 490), 9860
$ DL(7),DT(7),THETAH,DTIME,TIMFAC,TIMBGN, 9870
$ CEQUI,DFW,TIMSUM,IPRMH(3,7),MAT(300), 9880

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Table 2.--Computer-program listing--Continued

\$	NBP( 490),IELMR	9890
	COMMON/SINBLK/QEH(8),QEC(8),XORD( 490),YORD( 490),ZORD( 490),	9900
\$	QS( 980),CS( 980),QF(1),CF(1),NP(300,8),IBPRM(300)	9910
	COMMON/SENBLK/WH(30),	9920
\$	WC(30),THETAC,NPBC( 490,1),NODOBS( 490)	9930
	COMMON/LDUBLK/CPKC(490,180),CPKH(490, 90),SH(8, 490),	9940
\$	QC( 490),QH( 490),IBC,IBH	9950
	REAL*8 CPKC,CPKH,SH,QC,QH	9960
	COMMON/SHABLK/SF(4,8,8),WF(4,8),WT(8),NUMQPT	9970
	COMMON/LINBLK/RLF(2,2),RLWT,LINOD(4,2)	9980
	DIMENSION KARG(4)	9990
	DIMENSION DH(8,8),EH(8,8),RJAC(8,8),RJACI(8,8),	10000
\$	DNDX(8),DNDY(8),DNDZ(8)	10010
	DOUBLE PRECISION SF,WT,WF,QEH,QEC,DH,EH,RJAC,RJACI,DNDX,DNDY,	10020
\$	DNDZ,TEMP,TMP,TMP1,TMP2	10030
	DOUBLE PRECISION DETJ,RN,DXH,DYH,DZH,RNN	10040
	DO 260 I=1,NUMNP	10050
	N=NPBC(I,IBCNT)	10060
	I1=(IWCNT-1)*NUMNP+I	10070
	QH(I)=WELL(I1)	10080
	IF (N.GT.0) QH(I)=QH(I)*QF(N)	10090
	IF (NVARH.LE.0) GO TO 258	10100
	DO 255 J=1,NVARH	10110
	SH(J,I)=0.0	10120
255	CONTINUE	10130
258	N1=N+NPARG	10140
	IF (N1.GT.NPARH) SH(N1,I)=SH(N1,I)+QH(I)	10150
260	CONTINUE	10160
	IF (LU.EQ.0) GO TO 266	10170
	N=NUMNP+NOBMAX	10180
	DO 250 I=1,N	10190
	SFOBSH(I)=0.0	10200
250	CONTINUE	10210
	NCORNR=4	10220
	DO 265 I=1,IDIM	10230
	DO 262 J=1,IBH	10240
	CPKH(I,J)=0.0	10250
262	CONTINUE	10260
265	CONTINUE	10270
266	NODES=8	10280
	DO 399 I=1,NUMEL	10290
	MATI=IABS(MAT(I))	10300
	DO 305 J=1,NODES	10310
	QEH(J)=0.0	10320
	DO 305 K=1,NODES	10330
	DH(J,K)=0.0	10340
	EH(J,K)=0.0	10350
305	CONTINUE	10360
C**	VOLUME QUADRATURE	10370
	DO 340 J=1,NUMQPT	10380
	RJAC(1,1)=0.0	10390
	RJAC(1,2)=0.0	10400

Table 2.--Computer-program listing--Continued

```

RJAC(1,3)=0.0                                10410
RJAC(2,1)=0.0                                10420
RJAC(2,2)=0.0                                10430
RJAC(2,3)=0.0                                10440
RJAC(3,1)=0.0                                10450
RJAC(3,2)=0.0                                10460
RJAC(3,3)=0.0                                10470
308  L1=IPRMH(1,MATI)                          10480
      L2=IPRMH(2,MATI)                          10490
      L3=IPRMH(3,MATI)                          10500
      RXJ=RX(MATI)                              10510
      RYJ=RY(MATI)                              10520
      RZJ=RZ(MATI)                              10530
      RCJ=RC(MATI)                              10540
      DO 310 K=1,NODES                          10550
      NPK=NP(I,K)                              10560
      RJAC(1,1)=RJAC(1,1)+SF(2,K,J)*XORD(NPK)  10570
      RJAC(1,2)=RJAC(1,2)+SF(3,K,J)*XORD(NPK)  10580
      RJAC(1,3)=RJAC(1,3)+SF(4,K,J)*XORD(NPK)  10590
      RJAC(2,1)=RJAC(2,1)+SF(2,K,J)*YORD(NPK)  10600
      RJAC(2,2)=RJAC(2,2)+SF(3,K,J)*YORD(NPK)  10610
      RJAC(2,3)=RJAC(2,3)+SF(4,K,J)*YORD(NPK)  10620
      RJAC(3,1)=RJAC(3,1)+SF(2,K,J)*ZORD(NPK)  10630
      RJAC(3,2)=RJAC(3,2)+SF(3,K,J)*ZORD(NPK)  10640
      RJAC(3,3)=RJAC(3,3)+SF(4,K,J)*ZORD(NPK)  10650
310  CONTINUE                                  10660
      DETJ=RJAC(1,1)*(RJAC(2,2)*RJAC(3,3)-RJAC(3,2)*RJAC(2,3))  10670
      DETJ=-RJAC(1,2)*(RJAC(2,1)*RJAC(3,3)-RJAC(3,1)*RJAC(2,3))+DETJ  10680
      DETJ=RJAC(1,3)*(RJAC(2,1)*RJAC(3,2)-RJAC(3,1)*RJAC(2,2))+DETJ  10690
      IF (DETI.LE.0.0) GO TO 7001                10700
311  RJACI(1,1)=(RJAC(2,2)*RJAC(3,3)-RJAC(3,2)*RJAC(2,3))/DETI  10710
      RJACI(1,2)=-((RJAC(1,2)*RJAC(3,3)-RJAC(3,2)*RJAC(1,3))/DETI  10720
      RJACI(1,3)=(RJAC(1,2)*RJAC(2,3)-RJAC(2,2)*RJAC(1,3))/DETI  10730
      RJACI(2,1)=-((RJAC(2,1)*RJAC(3,3)-RJAC(3,1)*RJAC(2,3))/DETI  10740
      RJACI(2,2)=(RJAC(1,1)*RJAC(3,3)-RJAC(3,1)*RJAC(1,3))/DETI  10750
      RJACI(2,3)=-((RJAC(1,1)*RJAC(2,3)-RJAC(2,1)*RJAC(1,3))/DETI  10760
      RJACI(3,1)=(RJAC(2,1)*RJAC(3,2)-RJAC(3,1)*RJAC(2,2))/DETI  10770
      RJACI(3,2)=-((RJAC(1,1)*RJAC(3,2)-RJAC(3,1)*RJAC(1,2))/DETI  10780
      RJACI(3,3)=(RJAC(1,1)*RJAC(2,2)-RJAC(2,1)*RJAC(1,2))/DETI  10790
      DO 315 K=1,NODES                          10800
      DNDX(K)=RJACI(1,1)*SF(2,K,J)+RJACI(2,1)*SF(3,K,J)+RJACI(3,1)  10810
      &      *SF(4,K,J)                          10820
      DNDY(K)=RJACI(1,2)*SF(2,K,J)+RJACI(2,2)*SF(3,K,J)+RJACI(3,2)  10830
      &      *SF(4,K,J)                          10840
      DNDZ(K)=RJACI(1,3)*SF(2,K,J)+RJACI(2,3)*SF(3,K,J)+RJACI(3,3)  10850
      &      *SF(4,K,J)                          10860
315  CONTINUE                                  10870
C** LINE INTEGRATION TO FILL SFOBSH            10880
      IF (NVARH.LE.0.OR.LU.EQ.0) GO TO 314      10890
      DO 280 K=1,NCORNR                        10900
      NP1=LINOD(K,1)                            10910
      NP2=LINOD(K,2)                            10920

```

Table 2.--Computer-program listing--Continued

	NP1=NP(I,NP1)	10930
	NP2=NP(I,NP2)	10940
	IF (NODOBS(NP1).LE.0.OR.NODOBS(NP2).LE.0) GO TO 280	10950
	IF (NODOBS(NP1).NE.NODOBS(NP2)) GO TO 280	10960
	N=NUMNP+NODOBS(NP1)	10970
	DETL=0.0	10980
	DO 268 L=1,2	10990
	NPL=LINOD(K,L)	11000
	NPL=NP(I,NPL)	11010
	DETL=DETL+RLF(2,L)*ZORD(NPL)	11020
268	CONTINUE	11030
	DETL=ABS(DETL)*WTNOD(NP1)*WTNOD(NP2)	11040
	DO 270 L=1,2	11050
	NPL=LINOD(K,L)	11060
	NPL=NP(I,NPL)	11070
	NPK=2	11080
	IF (L.EQ.2) NPK=1	11090
	NPK=LINOD(K,NPK)	11100
	NPK=NP(I,NPK)	11110
	SFOBSH(NPL)=SFOBSH(NPL)+RLWT*RLF(1,L)*DETL*WTNOD(NPL)*	11120
	\$ (RXJ+RYJ)	11130
	SFOBSH(NPK)=SFOBSH(NPK)+RLWT*RLF(1,L)*DETL*(1.0-WTNOD(NPL))*	11140
	\$ (RXJ+RYJ)	11150
	SFOBSH(N)=SFOBSH(N)+RLWT*RLF(1,L)*DETL*(RXJ+RYJ)	11160
270	CONTINUE	11170
280	CONTINUE	11180
C**	END OF LINE INTEGRATION	11190
314	DO 330 K=1,NODES	11200
	NPK=NP(I,K)	11210
	I1=NPK+NUMNP	11220
	I2=NPK	11230
	I4=NPK+NUMNP	11240
	I5=I4-NUMNP	11250
229	DO 329 L=1,NODES	11260
317	NPL=NP(I,L)+NUMNP	11270
	NPLM1=NPL-NUMNP	11280
	J1=NP(I,L)+NUMNP	11290
	J2=J1-NUMNP	11300
	RNN=WT(J)*SF(1,K,J)*SF(1,L,J)*DETJ	11310
	DXH=WT(J)*DNDX(K)*DNDX(L)*DETJ	11320
	DYH=WT(J)*DNDY(K)*DNDY(L)*DETJ	11330
	DZH=WT(J)*DNDZ(K)*DNDZ(L)*DETJ	11340
	DH(K,L)=DH(K,L)+DXH*RXJ+DYH*RYJ+DZH*RZJ	11350
	EH(K,L)=EH(K,L)+RNN*RCJ	11360
	IF (L1.GT.0) SH(L1,NPK)=SH(L1,NPK)+DXH*(H(J1)*THETAH-H(J2)*(1.0-	11370
	\$ THETAH))*RXJ	11380
	IF (L2.GT.0) SH(L2,NPK)=SH(L2,NPK)+DYH	11390
	\$ *(H(J1)*THETAH-H(J2)*(1.0-THETAH))*RYJ	11400
	IF (L3.GT.0) SH(L3,NPK)=SH(L3,NPK)+DZH	11410
	\$ *(H(J1)*THETAH-H(J2)*(1.0-THETAH))*RZJ	11420
329	CONTINUE	11430
330	CONTINUE	11440

Table 2.--Computer-program listing--Continued

340	CONTINUE	11450
C**	SURFACE QUADRATURE	11460
	IF (MAT(I).GE.0) GO TO 382	11470
360	DO 369 J=1,6	11480
	GO TO (362,364,365,366,367,359),J	11490
362	K1=NP(I,1)+(ISCNT-1)*NUMNP	11500
	K2=NP(I,2)+(ISCNT-1)*NUMNP	11510
	K3=NP(I,3)+(ISCNT-1)*NUMNP	11520
	K4=NP(I,4)+(ISCNT-1)*NUMNP	11530
	NN=4	11540
	GO TO 368	11550
364	K1=NP(I,5)+(ISCNT-1)*NUMNP	11560
	K2=NP(I,6)+(ISCNT-1)*NUMNP	11570
	K3=NP(I,7)+(ISCNT-1)*NUMNP	11580
	K4=NP(I,8)+(ISCNT-1)*NUMNP	11590
	NN=4	11600
	GO TO 368	11610
365	K1=NP(I,1)+(ISCNT-1)*NUMNP	11620
	K2=NP(I,5)+(ISCNT-1)*NUMNP	11630
	K3=NP(I,6)+(ISCNT-1)*NUMNP	11640
	K4=NP(I,2)+(ISCNT-1)*NUMNP	11650
	NN=4	11660
	GO TO 368	11670
366	K1=NP(I,1)+(ISCNT-1)*NUMNP	11680
	K2=NP(I,4)+(ISCNT-1)*NUMNP	11690
	K3=NP(I,8)+(ISCNT-1)*NUMNP	11700
	K4=NP(I,5)+(ISCNT-1)*NUMNP	11710
	NN=4	11720
	GO TO 368	11730
367	K1=NP(I,2)+(ISCNT-1)*NUMNP	11740
	K2=NP(I,6)+(ISCNT-1)*NUMNP	11750
	K3=NP(I,7)+(ISCNT-1)*NUMNP	11760
	K4=NP(I,3)+(ISCNT-1)*NUMNP	11770
	NN=4	11780
	GO TO 368	11790
359	K1=NP(I,3)+(ISCNT-1)*NUMNP	11800
	K2=NP(I,4)+(ISCNT-1)*NUMNP	11810
	K3=NP(I,8)+(ISCNT-1)*NUMNP	11820
	K4=NP(I,7)+(ISCNT-1)*NUMNP	11830
	NN=4	11840
368	CHKQ=QS(K1)*QS(K2)*QS(K3)*QS(K4)	11850
	IF (CHKQ.EQ.0.0) GO TO 369	11860
	KARG(1)=K1	11870
	KARG(2)=K2	11880
	KARG(3)=K3	11890
	KARG(4)=K4	11900
	CALL SURINT(I,NODES,KARG,NN,NPARH,0,ISCNT,NUMNP)	11910
369	CONTINUE	11920
C**	PLACE IN LARGE CPKH AND Q MATRICES	11930
382	DO 386 J=1,NODES	11940
	NPJ=NP(I,J)	11950
	QH(NPJ)=QH(NPJ)+QE(J)	11960

Table 2.--Computer-program listing--Continued

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379 DO 384 K=1,NODES 11970
    NPK=NP(I,K) 11980
    K1=NPK-NPJ+1 11990
390 QH(NPJ)=QH(NPJ)+(EH(J,K)/DTIME/FLOAT(NTPER)- 12000
    $ (1.0-THETAH)*DH(J,K))*H(I2) 12010
389 IF (NPBC(NPJ,IBCNT).GE.0.OR.NPBC(NPJ,IBCNT).EQ.-2.OR. 12020
    $ NPBC(NPJ,IBCNT).EQ.-4) GO TO 388 12030
    IF (NBP(NPJ).LE.0) GO TO 388 12040
    IF (NBP(NPJ).EQ.NBP(NPK)) GO TO 388 12050
    N=NBP(NPJ)+NQPAR+NPARH 12060
    SH(N,NPK)=SH(N,NPK)+(EH(J,K)/DTIME/FLOAT(NTPER)+THETAH*D 12070
    $ H(NPJ+NUMNP)-(EH(J,K)/DTIME/FLOAT(NTPER)-(1.0- 12080
    $ THETAH)*DH(J,K))*H(NPJ) 12090
388 IF (NPK.LT.NPJ) GO TO 385 12100
    IF (LU.EQ.1) 12110
    $CPKH(NPJ,K1)=CPKH(NPJ,K1)+EH(J,K)/DTIME/FLOAT(NTPER)+THETAH* 12120
    $ DH(J,K) 12130
385 CONTINUE 12140
384 CONTINUE 12150
386 CONTINUE 12160
399 CONTINUE 12170
C** BOUNDARY CONDITIONS 12180
    DO 420 I=1,NUMNP 12190
    IF (NPBC(I,IBCNT).GE.0.OR.NPBC(I,IBCNT).EQ.-4) GO TO 420 12200
    IF (NPBC(I,IBCNT).EQ.-2) GO TO 420 12210
    IF (LU.EQ.1) CPKH(I,1)=1.0D+50 12220
    I1=I+NUMNP 12230
    QH(I)=H(I1)*1.0D+50 12240
420 CONTINUE 12250
    IF (LU.EQ.0.OR.NVARH.LE.0) RETURN 12260
C** FINISH COMPUTATION OF SFOBSH 12270
    DO 294 I=1,NUMNP 12280
    N=NUMNP+NODOBS(I) 12290
    IF (N.EQ.NUMNP) GO TO 294 12300
    SFOBSH(I)=SFOBSH(I)/SFOBSH(N) 12310
294 CONTINUE 12320
    RETURN 12330
7001 WRITE (6,17) I,J 12340
17 FORMAT ('OELEMENT',I5,' HAS NEGATIVE OR ZERO DETJ AT ', 12350
    $ 'QUADRATURE POINT',I5) 12360
    STOP 12370
    END 12380
    SUBROUTINE SALT (IFLOW,IWCNT,ISCNT,IBCNT,KT,KTH,NTPER,NUMNP, 12390
    $ LU,NUMEL,ISTEDY,IVELPR,IFRAC,IHCFLG, 12400
    $ ITER,IDIM,ALFAX,ALFAY,ALFAZ,ALFLG,NTFRAC,ITFRAC, 12410
    $ NOBMAX,IOBCNT) 12420
C** ROUTINE TO FORMULATE GLOBAL STIFFNESS AND SENSITIVITY MATRICES 12430
    COMMON/COFBLK/CFRAC(10,490),QHTMP(490),CINIT( 490), 12440
    $ H( 980),C( 980),WELL( 490),WELLC( 490), 12450
    $ RX(7),RY(7),RZ(7),RC(7), 12460
    $ RPR(7),DM(7),FP(7),FQ(7),RMT(7),HO(30), 12470
    $ RPBR(7),CBINIT(7),SFOBSH(520),SFOBSC(520),WTNOD( 490), 12480

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Table 2.--Computer-program listing--Continued

\$	DL(7),DT(7),THETAH,DTIME,TIMFAC,TIMBGN,	12490
\$	CEQUI,DFW,TIMSUM,IPRMH(3,7),MAT(300),NBP( 490),IELMR	12500
	COMMON/SINBLK/QEH(8),QEC(8),XORD( 490),YORD( 490),ZORD( 490),	12510
\$	QS( 980),CS( 980),QF(1),CF(1),NP(300,8),IBPRM(300)	12520
	COMMON/SENBLK/WH(30),	12530
\$	WC(30),THETAC,NPBC( 490,1),NODOBS( 490)	12540
	COMMON/LDUBLK/CPKC(490,180),CPKH(490, 90),SH(8, 490),	12550
\$	QC( 490),QH( 490),IBC,IBH	12560
	REAL*8 CPKC,CPKH,SH,QC,QH	12570
	COMMON/SHABLK/SF(4,8,8),WF(4,8),WT(8),NUMQPT	12580
	COMMON/LINBLK/RLF(2,2),RLWT,LINOD(4,2)	12590
	DIMENSION DC(8,8),EC(8,8),RJAC(8,8),RJACI(8,8),DNDX(8),DNDY(8),	12600
\$	DNDZ(8),KARG(4)	12610
	DOUBLE PRECISION SF,WT,WF,QEH,QEC,DC,EC,RJAC,RJACI,	12620
\$	DNDX,DNDY,DNDZ,TEMP,TMP,TMP1,TMP2,DFDX,	12630
\$	DFDY,DFDZ	12640
	DOUBLE PRECISION RN,DXH,DYH,DZH,RNN,DXYC,DXZC,DYZC,DNXN,DNYN,	12650
\$	DNZN,RNSUM	12660
	IDIAG=(IBC-1)/2+1	12670
	THETMP=(FLOAT(IFLOW-1)+THETAC)/FLOAT(NTPER)	12680
	IF (ISTEDY.EQ.1) THETMP=1.0	12690
	N=NUMNP+NOBMAX	12700
	DO 250 I=1,N	12710
	SFOBSC(I)=0.0	12720
250	CONTINUE	12730
	NCORNR=4	12740
	DO 260 I=1,NUMNP	12750
	N=NPBC(I,IBCNT)	12760
	I1=(IWCNT-1)*NUMNP+I	12770
241	QC(I)=0.0	12780
	IF (WELL(I1).LE.0.0) GO TO 260	12790
	QC(I)=WELL(I1)*WELLC(I1)	12800
	IF (N.GT.0) QC(I)=QC(I)*CF(N)	12810
260	CONTINUE	12820
	IF (LU.EQ.0) GO TO 266	12830
	DO 265 I=1,IDIM	12840
	DO 264 J=1,IBC	12850
	CPKC(I,J)=0.0	12860
264	CONTINUE	12870
265	CONTINUE	12880
266	IF (IVELPR.EQ.1) WRITE (6,18)	12890
18	FORMAT ('OELEMENT VX',11X,'VY',11X,'VZ',10X,'DISPX',8X,'DISPY',	12900
\$	8X,'DISPZ')	12910
	NODES=8	12920
	DO 399 I=1,NUMEL	12930
	MATI=IABS(MAT(I))	12940
	DO 305 J=1,NODES	12950
	QEC(J)=0.0	12960
	DO 305 K=1,NODES	12970
261	DC(J,K)=0.0	12980
	EC(J,K)=0.0	12990
305	CONTINUE	13000

Table 2.--Computer-program listing--Continued

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C** VOLUME QUADRATURE 13010
DO 340 J=1,NUMQPT 13020
XJ=0.0 13030
YJ=0.0 13040
ZJ=0.0 13050
VXJ=0.0 13060
VYJ=0.0 13070
VZJ=0.0 13080
RJAC(1,1)=0.0 13090
RJAC(1,2)=0.0 13100
RJAC(1,3)=0.0 13110
RJAC(2,1)=0.0 13120
RJAC(2,2)=0.0 13130
RJAC(2,3)=0.0 13140
RJAC(3,1)=0.0 13150
RJAC(3,2)=0.0 13160
RJAC(3,3)=0.0 13170
DO 310 K=1,NODES 13180
NPK=NP(I,K) 13190
RJAC(1,1)=RJAC(1,1)+SF(2,K,J)*XORD(NPK) 13200
RJAC(1,2)=RJAC(1,2)+SF(3,K,J)*XORD(NPK) 13210
RJAC(1,3)=RJAC(1,3)+SF(4,K,J)*XORD(NPK) 13220
RJAC(2,1)=RJAC(2,1)+SF(2,K,J)*YORD(NPK) 13230
RJAC(2,2)=RJAC(2,2)+SF(3,K,J)*YORD(NPK) 13240
RJAC(2,3)=RJAC(2,3)+SF(4,K,J)*YORD(NPK) 13250
RJAC(3,1)=RJAC(3,1)+SF(2,K,J)*ZORD(NPK) 13260
RJAC(3,2)=RJAC(3,2)+SF(3,K,J)*ZORD(NPK) 13270
RJAC(3,3)=RJAC(3,3)+SF(4,K,J)*ZORD(NPK) 13280
310 CONTINUE 13290
DETJ=RJAC(1,1)*(RJAC(2,2)*RJAC(3,3)-RJAC(3,2)*RJAC(2,3)) 13300
DETJ=-RJAC(1,2)*(RJAC(2,1)*RJAC(3,3)-RJAC(3,1)*RJAC(2,3))+DETJ 13310
DETJ=RJAC(1,3)*(RJAC(2,1)*RJAC(3,2)-RJAC(3,1)*RJAC(2,2))+DETJ 13320
IF (DETJ.LE.0.0) GO TO 7001 13330
RJACI(1,1)=(RJAC(2,2)*RJAC(3,3)-RJAC(3,2)*RJAC(2,3))/DETJ 13340
RJACI(1,2)=-(RJAC(1,2)*RJAC(3,3)-RJAC(3,2)*RJAC(1,3))/DETJ 13350
RJACI(1,3)=(RJAC(1,2)*RJAC(2,3)-RJAC(2,2)*RJAC(1,3))/DETJ 13360
RJACI(2,1)=-(RJAC(2,1)*RJAC(3,3)-RJAC(3,1)*RJAC(2,3))/DETJ 13370
RJACI(2,2)=(RJAC(1,1)*RJAC(3,3)-RJAC(3,1)*RJAC(1,3))/DETJ 13380
RJACI(2,3)=-(RJAC(1,1)*RJAC(2,3)-RJAC(2,1)*RJAC(1,3))/DETJ 13390
RJACI(3,1)=(RJAC(2,1)*RJAC(3,2)-RJAC(3,1)*RJAC(2,2))/DETJ 13400
RJACI(3,2)=-(RJAC(1,1)*RJAC(3,2)-RJAC(3,1)*RJAC(1,2))/DETJ 13410
RJACI(3,3)=(RJAC(1,1)*RJAC(2,2)-RJAC(2,1)*RJAC(1,2))/DETJ 13420
DO 315 K=1,NODES 13430
DNDX(K)=RJACI(1,1)*SF(2,K,J)+RJACI(2,1)*SF(3,K,J)+RJACI(3,1) 13440
& *SF(4,K,J) 13450
DNDY(K)=RJACI(1,2)*SF(2,K,J)+RJACI(2,2)*SF(3,K,J)+RJACI(3,2) 13460
& *SF(4,K,J) 13470
DNDZ(K)=RJACI(1,3)*SF(2,K,J)+RJACI(2,3)*SF(3,K,J)+RJACI(3,3) 13480
& *SF(4,K,J) 13490
315 CONTINUE 13500
DFDX=RJACI(1,1)*WF(2,J)+RJACI(2,1)*WF(3,J)+RJACI(3,1)*WF(4,J) 13510
DFDY=RJACI(1,2)*WF(2,J)+RJACI(2,2)*WF(3,J)+RJACI(3,2)*WF(4,J) 13520

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Table 2.--Computer-program listing--Continued

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      DFDZ=RJACI(1,3)*WF(2,J)+RJACI(2,3)*WF(3,J)+RJACI(3,3)*WF(4,J)      13530
C** VELOCITY AND DISPERSION CALCULATIONS      13540
308  L1=IPRMH(1,MATI)      13550
      L2=IPRMH(2,MATI)      13560
      L3=IPRMH(3,MATI)      13570
      RXJ=RX(MATI)      13580
      RYJ=RY(MATI)      13590
      RZJ=RZ(MATI)      13600
      RCJ=RC(MATI)      13610
      PORBJ=RPBR(MATI)      13620
      PORJ=RPR(MATI)      13630
      RMASSJ=RMT(MATI)      13640
      DO 311 K=1,NODES      13650
      J1=NP(I,K)+NUMNP      13660
      J2=J1-NUMNP      13670
      VXJ=VXJ-DNDX(K)*RXJ*(THETMP*H(J1)+(1.0-THETMP)*H(J2))      13680
      VYJ=VYJ-DNDY(K)*RYJ*(THETMP*H(J1)+(1.0-THETMP)*H(J2))      13690
      VZJ=VZJ-DNDZ(K)*RZJ*(THETMP*H(J1)+(1.0-THETMP)*H(J2))      13700
311  CONTINUE      13710
      VMAG=(VXJ*VXJ+VYJ*VYJ+VZJ*VZJ)**0.5      13720
      IF (ABS(VMAG).LT.1.0E-20) GO TO 313      13730
312  DSPXJ=(DL(MATI)*VXJ*VXJ+DT(MATI)*VYJ*VYJ+DT(MATI)*VZJ*VZJ)/VMAG      13740
      $      +DFW      13750
      DSPYJ=(DT(MATI)*VXJ*VXJ+DL(MATI)*VYJ*VYJ+DT(MATI)*VZJ*VZJ)/VMAG      13760
      $      +DFW      13770
      DSPZJ=(DT(MATI)*VXJ*VXJ+DT(MATI)*VYJ*VYJ+DL(MATI)*VZJ*VZJ)/VMAG      13780
      $      +DFW      13790
301  DSPXYJ=(DL(MATI)-DT(MATI))*VXJ*VYJ/VMAG+DFW      13800
      DSPXZJ=(DL(MATI)-DT(MATI))*VXJ*VZJ/VMAG+DFW      13810
      DSPYZJ=(DL(MATI)-DT(MATI))*VYJ*VZJ/VMAG+DFW      13820
313  IF (IVELPR.EQ.1.AND.J.EQ.1) WRITE (6,19) I,VXJ,VYJ,VZJ,DSPXJ,      13830
      $      DSPYJ,DSPZJ      13840
19   FORMAT (I5,6(1X,E12.5))      13850
C** LINE INTEGRATION TO FILL SFOBSC      13860
      IF (NOBMAX.LE.0) GO TO 314      13870
      DO 280 K=1,NCORNR      13880
      NP1=LINOD(K,1)      13890
      NP2=LINOD(K,2)      13900
      NP1=NP(I,NP1)      13910
      NP2=NP(I,NP2)      13920
      IF (NODOBS(NP1).LE.0.OR.NODOBS(NP2).LE.0) GO TO 280      13930
      IF (NODOBS(NP1).NE.NODOBS(NP2)) GO TO 280      13940
      N=NUMNP+NODOBS(NP1)      13950
      DETL=0.0      13960
      DO 268 L=1,2      13970
      NPL=LINOD(K,L)      13980
      NPL=NP(I,NPL)      13990
      DETL=DETL+RLF(2,L)*ZORD(NPL)      14000
268  CONTINUE      14010
      DETL=ABS(DETL)*WTNOD(NP1)*WTNOD(NP2)      14020
C** ADJUST WEIGHTS FOR OBSERVED HEAD .GT. COMPUTED HEAD      14030
      N1=N-NUMNP+NOBMAX*(IOBCNT-1)      14040

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Table 2.--Computer-program listing--Continued

	HOBS=HO(N1)	14050
	H1=H(NP1+NUMNP)	14060
	H2=H(NP2+NUMNP)	14070
	HWT1=1.0	14080
	HWT2=1.0	14090
	IF (H1.GE.HOBS.AND.H2.GE.HOBS) GO TO 269	14100
	IF (H1.GE.HOBS.AND.H2.LT.HOBS) HWT2=(HOBS-H1)/(H2-H1)	14110
	IF (H1.LT.HOBS.AND.H2.GE.HOBS) HWT1=(HOBS-H2)/(H1-H2)	14120
	IF (HWT1.GE.1.0.AND.HWT2.GE.1.0) GO TO 314	14130
269	DO 270 L=1,2	14140
	NPL=LINOD(K,L)	14150
	NPL=NP(I,NPL)	14160
	NPK=2	14170
	IF (L.EQ.1) GO TO 272	14180
	NPK=1	14190
	HWT1=HWT2	14200
	H1=H2	14210
272	NPK=LINOD(K,NPK)	14220
	NPK=NP(I,NPK)	14230
	SFOBSC(NPL)=SFOBSC(NPL)+RLWT*RLF(1,L)*DETL*WTNOD(NPL)*HWT1*	14240
	\$ (RXJ+RYJ)*(H1-HOBS)	14250
	SFOBSC(NPK)=SFOBSC(NPK)+RLWT*RLF(1,L)*DETL*(1.0-WTNOD(NPL))*HWT1*	14260
	\$ (RXJ+RYJ)*(H1-HOBS)	14270
	SFOBSC(N)=SFOBSC(N)+RLWT*RLF(1,L)*DETL*HWT1*	14280
	\$ (RXJ+RYJ)*(H1-HOBS)	14290
270	CONTINUE	14300
280	CONTINUE	14310
C**	END OF LINE INTEGRATION	14320
314	DO 330 K=1,NODES	14330
	NPK=NP(I,K)	14340
	I1=NPK+NUMNP	14350
	I2=NPK	14360
	I4=NPK+NUMNP	14370
	I5=I4-NUMNP	14380
319	TMP=WT(J)*SF(1,K,J)*RMASSJ*CEQUI*DETJ	14390
	QEC(K)=QEC(K)+TMP	14400
322	CONTINUE	14410
C**	FRACTURE SOURCE-SINK TERM	14420
100	IF (IFRAC.EQ.0.OR.PORBJ.EQ.0.0) GO TO 199	14430
	PFRACJ=FP(MATI)	14440
	QBLOJ=FQ(MATI)	14450
	PORJ=RPR(MATI)	14460
	CB=0.0	14470
	DTMP=DTIME	14480
	TMPSUM=TIMSUM	14490
	PI2=9.869604404	14500
	ITI=NTFRAC+2	14510
	I1=NTFRAC	14520
	IF (KT.LT.NTFRAC) I1=KT	14530
	IF (I1.LT.2) GO TO 105	14540
C	TIME STEPS 1 THROUGH M-1	14550
	DO 103 IF1=1,I1	14560

Table 2.--Computer-program listing--Continued

	ITI=ITI-1	14570
	ITIM1=ITI-1	14580
	TMPSUM=TMPSUM-DTMP	14590
	DTMP=DTMP/TIMFAC	14600
	TMP=0.0	14610
	DO 101 IF2=1,ITFRAC	14620
	EN=(2.0*(IF2-1)+1)**2	14630
	BN=DM(MATI)*EN*PI2/QBLOCJ**2	14640
	TMP1=(EXP(-BN*(TIMSUM-TMPSUM))-EXP(-BN*(TIMSUM-TMPSUM+DTMP)))	14650
	\$ /EN/BN	14660
	TMP=TMP-TMP1	14670
	TMP1=(EXP(-BN*(TIMSUM-DTIME-TMPSUM))-EXP(-BN*(TIMSUM-DTIME-TMPSUM	14680
	\$ +DTMP)))/EN/BN	14690
	TMP=TMP+TMP1	14700
101	CONTINUE	14710
	C1=CFRAC(ITI,NPK)	14720
	C0=CFRAC(ITIM1,NPK)	14730
	CB=CB+(C1-C0)*8.0*TMP/PI2/DTMP	14740
103	CONTINUE	14750
C	INITIAL TERM	14760
105	TMP=0.0	14770
	DO 107 IF2=1,ITFRAC	14780
	EN=(2.0*(IF2-1)+1)**2	14790
	BN=DM(MATI)*EN*PI2/QBLOCJ**2	14800
	TMP1=EXP(-BN*TIMSUM)/EN	14810
	TMP=TMP-TMP1	14820
	TMP1=EXP(-BN*(TIMSUM-DTIME))/EN	14830
	TMP=TMP+TMP1	14840
107	CONTINUE	14850
	C0=CINIT(NPK)	14860
	CB=CB+C0*8.0*TMP/PI2	14870
C	TIME STEP M	14880
110	NTF1=NTFRAC+1	14890
	TMP=0.0	14900
	DO 111 IF2=1,ITFRAC	14910
	EN=(2.0*(IF2-1)+1)**2	14920
	BN=DM(MATI)*EN*PI2/QBLOCJ**2	14930
	TMP=TMP+(1.0-EXP(-BN*DTIME))/EN/BN	14940
111	CONTINUE	14950
	CB=CB-(1.0-8.0*TMP/PI2/DTIME)*CFRAC(NTF1,NPK)	14960
	DCFRA=PORBJ*QBLOCJ*(1.0-8.0*TMP/DTIME/PI2)/DTIME/(PFRACJ+QBLOCJ)	14970
109	QXYFRA=-PORBJ*QBLOCJ*CB/DTIME/(PFRACJ+QBLOCJ)	14980
190	RN=WT(J)*SF(1,K,J)*DETJ	14990
	RNN=WT(J)*SF(1,K,J)*SF(1,K,J)*DETJ	15000
	QEC(K)=QEC(K)+RN*QXYFRA	15010
	DC(K,K)=DC(K,K)+RNN*NODES*DCFRA	15020
199	CONTINUE	15030
C**	FILL ELEMENT MATRICES	15040
	DO 329 L=1,NODES	15050
317	NPL=NP(I,L)+NUMNP	15060
	NPLM1=NPL-NUMNP	15070
	J1=NP(I,L)+NUMNP	15080

Table 2.--Computer-program listing--Continued

	J2=J1-NUMNP	15090
	RNN=WT(J)*SF(1,K,J)*SF(1,L,J)*DETJ	15100
	DXH=WT(J)*DETJ*DNDX(K)*DNDX(L)	15110
	DYH=WT(J)*DETJ*DNDY(K)*DNDY(L)	15120
	DZH=WT(J)*DETJ*DNDZ(K)*DNDZ(L)	15130
	DXYC=WT(J)*DETJ*DNDX(K)*DNDY(L)	15140
	DXZC=WT(J)*DETJ*DNDX(K)*DNDZ(L)	15150
	DYZC=WT(J)*DETJ*DNDY(K)*DNDZ(L)	15160
	BSIGN=1.0	15170
	KB=K	15180
	LB=L	15190
	IF (NPBC(K,IBCNT).LT.-4.OR.NPBC(L,IBCNT).LT.-4) GO TO 225	15200
	BSIGN=-1.0	15210
	KB=L	15220
	LB=K	15230
225	DNXN=WT(J)*DETJ*DNDX(KB)*SF(1,LB,J)	15240
	DNYN=WT(J)*DETJ*DNDY(KB)*SF(1,LB,J)	15250
	DNZN=WT(J)*DETJ*DNDZ(KB)*SF(1,LB,J)	15260
	RNLB=WT(J)*DETJ*SF(1,LB,J)	15270
	RNKB=WT(J)*DETJ*SF(1,KB,J)	15280
318	IF (ALFLG.LE.0.0) GO TO 150	15290
	ALFAX=ABS(ALFAX)	15300
	IF (VXJ.LT.0.0) ALFAX=-ALFAX	15310
	ALFAY=ABS(ALFAY)	15320
	IF (VYJ.LT.0.0) ALFAY=-ALFAY	15330
	ALFAZ=ABS(ALFAZ)	15340
	IF (VZJ.LT.0.0) ALFAZ=-ALFAZ	15350
	GO TO 197	15360
150	PECNUM=ABS(VXJ*(XORD(NPL)-XORD(NPK))/DSPXJ)	15370
	ALFAX=0.0	15380
	IF (PECNUM.GE.2.0) ALFAX=1.0/TANH(PECNUM/2.0)-2.0/PECNUM	15390
	IF (VXJ.LT.0.0) ALFAX=-ALFAX	15400
	PECNUM=ABS(VYJ*(YORD(NPL)-YORD(NPK))/DSPYJ)	15410
	ALFAY=0.0	15420
	IF (PECNUM.GE.2.0) ALFAY=1.0/TANH(PECNUM/2.0)-2.0/PECNUM	15430
	IF (VYJ.LT.0.0) ALFAY=-ALFAY	15440
	PECNUM=ABS(VZJ*(ZORD(NPL)-ZORD(NPK))/DSPZJ)	15450
	ALFAZ=0.0	15460
	IF (PECNUM.GE.2.0) ALFAZ=1.0/TANH(PECNUM/2.0)-2.0/PECNUM	15470
	IF (VZJ.LT.0.0) ALFAZ=-ALFAZ	15480
197	TMP=(DXH+ALFAX*DFDX*DNDX(L))*DSPXJ+(DYH+ALFAY*DFDY*DNDY(L))*DSPYJ+	15490
	\$ (DZH+ALFAZ*DFDZ*DNDZ(L))*DSPZJ+(DXZC+ALFAX*DFDX*DNDZ(L))*	15500
	\$ DSPXZJ+(DYZC+ALFAY*DFDY*DNDZ(L))*DSPYZJ+(DXYC+ALFAX*DFDX*	15510
	\$ DNDY(L))*DSPXYJ+(DXZC+ALFAZ*DFDZ*DNDX(L))*DSPXZJ+(DYZC+	15520
	\$ ALFAZ*DFDZ*DNDY(L))*DSPYZJ+(DXYC+ALFAY*DFDY*DNDX(L))*DSPXYJ	15530
	TEMP=((DNXN+ALFAX*DFDX*RNLB)*VXJ+(DNYN+ALFAY*DFDY*RNLB)*VYJ+	15540
	\$ (DNZN+ALFAZ*DFDZ*RNLB)*VZJ)*BSIGN	15550
	DC(K,L)=DC(K,L)+(TMP-TEMP)	15560
	EC(K,L)=EC(K,L)+RNN*PORJ+RNN*RMASJ	15570
329	CONTINUE	15580
330	CONTINUE	15590
340	CONTINUE	15600

Table 2.--Computer-program listing--Continued

C** SURFACE QUADRATURE	15610
IF (MAT(I).GE.0) GO TO 382	15620
360 DO 369 J=1,6	15630
GO TO (362,364,365,366,367,359),J	15640
362 K1=NP(I,1)+(ISCNT-1)*NUMNP	15650
K2=NP(I,2)+(ISCNT-1)*NUMNP	15660
K3=NP(I,3)+(ISCNT-1)*NUMNP	15670
K4=NP(I,4)+(ISCNT-1)*NUMNP	15680
NN=4	15690
GO TO 368	15700
364 K1=NP(I,5)+(ISCNT-1)*NUMNP	15710
K2=NP(I,6)+(ISCNT-1)*NUMNP	15720
K3=NP(I,7)+(ISCNT-1)*NUMNP	15730
K4=NP(I,8)+(ISCNT-1)*NUMNP	15740
NN=4	15750
GO TO 368	15760
365 K1=NP(I,1)+(ISCNT-1)*NUMNP	15770
K2=NP(I,5)+(ISCNT-1)*NUMNP	15780
K3=NP(I,6)+(ISCNT-1)*NUMNP	15790
K4=NP(I,2)+(ISCNT-1)*NUMNP	15800
NN=4	15810
GO TO 368	15820
366 K1=NP(I,1)+(ISCNT-1)*NUMNP	15830
K2=NP(I,4)+(ISCNT-1)*NUMNP	15840
K3=NP(I,8)+(ISCNT-1)*NUMNP	15850
K4=NP(I,5)+(ISCNT-1)*NUMNP	15860
NN=4	15870
GO TO 368	15880
367 K1=NP(I,2)+(ISCNT-1)*NUMNP	15890
K2=NP(I,6)+(ISCNT-1)*NUMNP	15900
K3=NP(I,7)+(ISCNT-1)*NUMNP	15910
K4=NP(I,3)+(ISCNT-1)*NUMNP	15920
NN=4	15930
GO TO 368	15940
359 K1=NP(I,3)+(ISCNT-1)*NUMNP	15950
K2=NP(I,4)+(ISCNT-1)*NUMNP	15960
K3=NP(I,8)+(ISCNT-1)*NUMNP	15970
K4=NP(I,7)+(ISCNT-1)*NUMNP	15980
NN=4	15990
368 CHKQ=QS(K1)*QS(K2)*QS(K3)*QS(K4)	16000
IF (CHKQ.EQ.0.0) GO TO 369	16010
KARG(1)=K1	16020
KARG(2)=K2	16030
KARG(3)=K3	16040
KARG(4)=K4	16050
CALL SURINT(I,NODES,KARG,NN,NPARH,1,ISCNT,NUMNP)	16060
369 CONTINUE	16070
C** LUMP CAPACITY MATRIX ON DIAGONAL	16080
DO 376 J=1,NODES	16090
DO 376 K=1,NODES	16100
IF (J.EQ.K) EC(J,K)=EC(J,K)*8.0	16110
IF (J.NE.K) EC(J,K)=0.0	16120

Table 2.--Computer-program listing--Continued

376	CONTINUE	16130
C**	PLACE IN LARGE CPKC AND Q MATRICES	16140
382	DO 386 J=1,NODES	16150
	NPJ=NP(I,J)	16160
	QC(NPJ)=QC(NPJ)+QEC(J)	16170
	DO 384 K=1,NODES	16180
	NPK=NP(I,K)	16190
	K2=NPK-NPJ+IDIAG	16200
	QC(NPJ)=QC(NPJ)+(EC(J,K)/DTIME-(1.0-THETAC)*DC(J,K))*C(NPK)	16210
383	IF (LU.EQ.1) CPKC(NPJ,K2)=CPKC(NPJ,K2)+EC(J,K)/DTIME	16220
	\$ +THETAC*DC(J,K)	16230
384	CONTINUE	16240
386	CONTINUE	16250
399	CONTINUE	16260
C**	BOUNDARY CONDITIONS	16270
	DO 420 I=1,NUMNP	16280
	IF (NPBC(I,IBCNT).GE.-1.OR.NPBC(I,IBCNT).LE.-4) GO TO 420	16290
	IF (LU.EQ.1) CPKC(I,IDIAG)=1.0D+50	16300
	I1=I+NUMNP	16310
	QC(I)=C(I1)*1.0D+50	16320
420	CONTINUE	16330
	IF (NOBMAX.LE.0) RETURN	16340
C**	FINISH COMPUTATION OF SFOBSC	16350
	DO 294 I=1,NUMNP	16360
	N=NUMNP+NODOBS(I)	16370
	IF (N.EQ.NUMNP) GO TO 294	16380
	SFOBSC(I)=SFOBSC(I)/SFOBSC(N)	16390
294	CONTINUE	16400
	RETURN	16410
7001	WRITE (6,17) I,J	16420
17	FORMAT ('OELEMENT',I5,' HAS NEGATIVE OR ZERO DETJ AT ',	16430
	\$ 'QUADRATURE POINT',I5)	16440
	STOP	16450
	END	16460
C**	PERFORM SURFACE INTEGRATION	16470
	SUBROUTINE SURINT(I,N,L,N1,NPARH,IHCFLG,ISCNT,NUMNP)	16480
	COMMON/SINBLK/QEH(8),QEC(8),XORD( 490),YORD( 490),ZORD( 490),	16490
	\$ QS( 980),CS( 980),QF(1),CF(1),NP(300,8),IBPRM(300)	16500
	COMMON/SENBLK/WH(30),	16510
	\$ WC(30),THETAC,NPBC( 490,1),NODOBS( 490)	16520
	COMMON/SURBLK/AF(3,4),AWT,NQPTA	16530
	COMMON/LDUBLK/CPKC(490,180),CPKH(490, 90),SH(8, 490),	16540
	\$ QC( 490),QH( 490),IBC,IBH	16550
	REAL*8 CPKC,CPKH,SH,QC,QH	16560
	DOUBLE PRECISION QEH,QEC,AF,AWT,TEMP,RJAC,DETJ,TMP	16570
	DIMENSION RJAC(3,2),L(4)	16580
	N3=IBPRM(I)	16590
	N2=N3+NPARH	16600
	RJAC(1,1)=0.0	16610
	RJAC(1,2)=0.0	16620
	RJAC(2,1)=0.0	16630
	RJAC(2,2)=0.0	16640



Table 2.--Computer-program listing--Continued

```

RJAC(3,1)=0.0                                16650
RJAC(3,2)=0.0                                16660
QSJ=0.0                                       16670
CSJ=0.0                                       16680
DO 310 K=1,N1                                16690
N4=L(K)                                       16700
NPK=N4-(ISCNT-1)*NUMNP                       16710
RJAC(1,1)=RJAC(1,1)+AF(2,K)*XORD(NPK)        16720
RJAC(1,2)=RJAC(1,2)+AF(3,K)*XORD(NPK)        16730
RJAC(2,1)=RJAC(2,1)+AF(2,K)*YORD(NPK)        16740
RJAC(2,2)=RJAC(2,2)+AF(3,K)*YORD(NPK)        16750
RJAC(3,1)=RJAC(3,1)+AF(2,K)*ZORD(NPK)        16760
RJAC(3,2)=RJAC(3,2)+AF(3,K)*ZORD(NPK)        16770
TMP=QS(N4)                                   16780
IF (IHCFLG.EQ.1) TEMP=CS(N4)                 16790
IF (N3.GT.0) TMP=TMP*QF(N3)                  16800
IF (N3.GT.0.AND.IHCFLG.EQ.1) TEMP=TEMP*CF(N3) 16810
QSJ=QSJ+AF(1,K)*TMP                          16820
IF (IHCFLG.EQ.1) CSJ=CSJ+AF(1,K)*TMP         16830
310  CONTINUE                                16840
      DETJ=RJAC(1,1)*(RJAC(2,2)-RJAC(3,2))    16850
$      -RJAC(2,1)*(RJAC(1,2)-RJAC(3,2))      16860
$      +RJAC(3,1)*(RJAC(1,2)-RJAC(2,2))      16870
      IF (DETI.LE.0.0) DETJ=-DETI             16880
324  DO 330 K=1,N1                            16890
      DO 327 II=1,N                            16900
      IF (NP(I,II).NE.L(K)-(ISCNT-1)*NUMNP) GO TO 327 16910
      K1=II                                    16920
      GO TO 328                                16930
327  CONTINUE                                16940
328  NPK=L(K)-(ISCNT-1)*NUMNP                 16950
      TMP=AWT*AF(1,K)*QSJ*DETI                16960
      IF (IHCFLG.EQ.1) GO TO 326              16970
      IF (N2.GT.NPARH) SH(N2,NPK)=SH(N2,NPK)+TMP 16980
      QEH(K1)=QEH(K1)+TMP                     16990
      GO TO 330                                17000
326  IF (TMP.LT.0.0) GO TO 330                17010
      QEC(K1)=QEC(K1)+CSJ*TMP                 17020
330  CONTINUE                                17030
      RETURN                                  17040
      END                                    17050
      SUBROUTINE SENS (NVARH,NUMNP,IPO,KT,KTH,IOBCNT,
$      IXTRFL,NOBMAX,IBCNT)                  17060
C** SUBROUTINE TO SETUP AND SOLVE FOR SCALED SENSITIVITIES 17080
      COMMON/SENBLK/WH(30),                   17090
$      WC(30),THETAC,NPBC( 490,1),NODOBS( 490) 17100
      COMMON/LDUBLK/CPKC(490,180),CPKH(490, 90),SH(8, 490), 17110
$      QC( 490),QH( 490),IBC,IBH             17120
      REAL*8 CPKC,CPKH,SH,QC,QH              17130
      IF (NVARH.LE.0) RETURN                  17140
C** SOLVE FOR FLOW SENSITIVITIES              17150
      DO 224 I=1,NVARH                       17160

```

Table 2.--Computer-program listing--Continued

	DO 224 J=1,NUMNP	17170
	IF (NPBC(J,IBCNT).EQ.-1) SH(I,J)=0.0	17180
	IF (NPBC(J,IBCNT).EQ.-3) SH(I,J)=0.0	17190
	IF (NPBC(J,IBCNT).EQ.-5) SH(I,J)=0.0	17200
224	CONTINUE	17210
	NEM1=NUMNP-1	17220
	NUMEQ=NUMNP	17230
	DO 350 I=1,NEM1	17240
	JEND=NUMEQ-I+1	17250
	IF(JEND.GT.IBH) JEND=IBH	17260
	DO 340 J=2,JEND	17270
	J1=I+J-1	17280
	QH(J1)=QH(J1)-QH(I)*CPKH(I,J)	17290
	DO 338 K=1,NVARH	17300
	SH(K,J1)=SH(K,J1)-SH(K,I)*CPKH(I,J)	17310
338	CONTINUE	17320
340	CONTINUE	17330
350	CONTINUE	17340
C**	BACK SUBSTITUTION	17350
	QH(NUMEQ)=QH(NUMEQ)/CPKH(NUMEQ,1)	17360
	DO 352 K=1,NVARH	17370
	SH(K,NUMEQ)=SH(K,NUMEQ)/CPKH(NUMEQ,1)	17380
352	CONTINUE	17390
353	DO 370 I=1,NEM1	17400
	I1=NUMEQ-I	17410
	QH(I1)=QH(I1)/CPKH(I1,1)	17420
	DO 354 K=1,NVARH	17430
	SH(K,I1)=SH(K,I1)/CPKH(I1,1)	17440
354	CONTINUE	17450
355	JEND=NUMEQ-I1+1	17460
	IF(JEND.GT.IBH) JEND=IBH	17470
	DO 360 J=2,JEND	17480
	J1=I1+J-1	17490
	QH(I1)=QH(I1)-CPKH(I1,J)*QH(J1)	17500
	DO 356 K=1,NVARH	17510
	SH(K,I1)=SH(K,I1)-CPKH(I1,J)*SH(K,J1)	17520
356	CONTINUE	17530
360	CONTINUE	17540
370	CONTINUE	17550
	IF (IPO.NE.1) GO TO 232	17560
	WRITE (6,1) KT	17570
	DO 230 I=1,NUMNP	17580
	J=NODOBS(I)	17590
	IF (J.LE.0) GO TO 230	17600
	J=NOBMAX*(IOBCNT-1)+J	17610
	WRITE (6, 2)I,(SH(K,I),K=1,NVARH)	17620
230	CONTINUE	17630
232	CONTINUE	17640
	RETURN	17650
1	FORMAT (' SCALED FLOW SENSITIVITIES - TIME STEP ',I5/' NODE',	17660
\$	20X,'PARAMETER')	17670
2	FORMAT (1X,I5,7(1X,G11.5)/6X,7(1X,G11.5))	17680

Table 2.--Computer-program listing--Continued

END	17690
SUBROUTINE LSTSQ (C,B,RK,P,YSQ,AP,AMP,RP,RPF,NVAR,NVX2,NVX3,IPO,	17700
\$                  INDT,KOUNT,NVE,LDIM)	17710
DIMENSION P(NVE),RK(NVE),B(LDIM),C(NVE,NVE)	17720
C** CHECK FOR NONZERO MARQUARDT PARAMETER	17730
IF(AMP.LT.-.5) GO TO 105	17740
74 IF (NVAR.EQ.1) GO TO 190	17750
DO 80 I=1,NVAR	17760
TEMP=C(I,I)+RK(I)	17770
IF(TEMP.GT.1.E-10) GO TO 78	17780
WRITE(6,260) I	17790
INDT=1	17800
GO TO 80	17810
78 C(I,I)=TEMP*.5	17820
80 CONTINUE	17830
IF(INDT.GT.0) RETURN	17840
NM1=NVAR-1	17850
DO 100 J=1,NM1	17860
TEMP=C(J,J)	17870
JP1=J+1	17880
DO 90 I=JP1,NVAR	17890
C(I,J)=C(I,J)/(C(I,I)*TEMP)	17900
90 C(J,I)=C(I,J)	17910
B(J+NVX2)=(B(J+NVX2)+RK(J)*(P(J)-1.))/TEMP+RP*TEMP*(RPF*P(J)-1.)	17920
B(J)=B(J+NVX2)	17930
B(J+NVX3)=TEMP	17940
100 C(J,J)=1.+RP+AMP	17950
TEMP=C(NVAR,NVAR)	17960
B(NVX3)=(B(NVX3)+RK(NVAR)*(P(NVAR)-1.))/TEMP	17970
1+RP*TEMP*(RPF*P(NVAR)-1.)	17980
B(NVAR)=B(NVX3)	17990
B(NVAR+NVX3)=TEMP	18000
C(NVAR,NVAR)=1.+RP+AMP	18010
IF(IPO.NE.1) GO TO 105	18020
WRITE(6,250)	18030
CALL PRTOT(C,NVE,0,NVAR)	18040
WRITE(6,230) (B(I+NVX2),I=1,NVAR)	18050
C** SOLVE FOR B USING LDU FACTORIZATION:	18060
C** DECOMPOSITION AND FORWARD SUBSTITUTION	18070
105 DET=1.	18080
DO 140 K=1,NM1	18090
PIV=C(K,K)	18100
DET=DET*PIV	18110
IF(ABS(PIV).GT.1.E-10) GO TO 110	18120
WRITE(6,210)	18130
INDT=1	18140
RETURN	18150
110 PIV=1./PIV	18160
KP1=K+1	18170
DO 130 J=KP1,NVAR	18180
TMP=C(J,K)*PIV	18190
DO 120 I=J,NVAR	18200

Table 2.--Computer-program listing--Continued

120	C(I,J)=C(I,J)-TMP*C(I,K)	18210
130	B(J)=B(J)-TMP*B(K)	18220
	C(K,K)=PIV	18230
140	CONTINUE	18240
	DET=DET*C(NVAR,NVAR)	18250
	IF(ABS(C(NVAR,NVAR)).GT.1.E-10) GO TO 150	18260
	WRITE(6,210)	18270
	INDT=1	18280
	RETURN	18290
150	IF(AMP.LT.-.5) RETURN	18300
C**	BACK SUBSTITUTION	18310
	B(NVX2)=B(NVAR)/C(NVAR,NVAR)	18320
	B(NVAR)=B(NVX2)/B(NVAR+NVX3)	18330
	I=NVAR	18340
160	I=I-1	18350
	IF (I.LE.0) GO TO 200	18360
	IP1=I+1	18370
	SUM=0.	18380
	DO 170 J=IP1,NVAR	18390
170	SUM=SUM+C(J,I)*B(J+NVAR)	18400
	B(I+NVAR)=(B(I)-SUM)*C(I,I)	18410
	B(I)=B(I+NVAR)/B(I+NVX3)	18420
	GO TO 160	18430
C**	SOLUTION WHEN NVAR=1	18440
190	TEMP=C(1,1)+RK(1)	18450
	IF(TEMP.GT.1.E-10) GO TO 195	18460
	I=1	18470
	WRITE(6,260) I	18480
	INDT=1	18490
	RETURN	18500
195	B(4)=TEMP**.5	18510
	B(3)=(B(3)+RK(1)*(P(1)-1.))/B(4)+RP*B(4)*(RPF*P(1)-1.)	18520
	C(1,1)=1.+RP+AMP	18530
	DET=C(1,1)	18540
	B(2)=B(3)/C(1,1)	18550
	B(1)=B(2)/B(4)	18560
C**	ADJUST AND PRINT REGRESSION COEFFICIENTS	18570
200	WRITE(6,220) KOUNT,YSQ,DET	18580
	DO 202 J=1,NVAR	18590
202	B(J)=AP*B(J)	18600
	WRITE(6,230) (B(J),J=1,NVAR)	18610
	RETURN	18620
210	FORMAT (43HOLEAST SQUARES COEFFICIENT MATRIX SINGULAR;/35H SOLUTIO	18630
	1N FOR PARAMETERS NOT UNIQUE)	18640
220	FORMAT (1H0,14HITERATION NO. ,I3/1H ,6HYSQ = ,E12.5	18650
	1,2X,9HDET(C) = ,E12.5/1H ,24HREGRESSION COEFFICIENTS:)	18660
230	FORMAT ((1H ,8(E12.5,2X)))	18670
250	FORMAT(49H0 SCALED LEAST SQUARES MATRIX AND GRADIENT VECTOR)	18680
260	FORMAT (29H0SENSITIVITIES FOR PARAMETER ,I4,17H EFFECTIVELY ZERO)	18690
	END	18700
	SUBROUTINE PRTOT(C,NO,IT,NUM)	18710
C**	IF IT=0, PRINT MATRICES DIVIDED VERTICALLY INTO BETWEEN ONE AND	18720

Table 2.--Computer-program listing--Continued

C**	TEN PARTS	18730
C**	IF IT=1, PRINT VECTOR IN THREE COLUMNS	18740
	DIMENSION C(1)	18750
	IF(IT.EQ.1) GO TO 25	18760
	ITMP=(NUM-1)/10+1	18770
	IB=1	18780
	DO 20 IBK=1,ITMP	18790
	INC=IBK*10	18800
	IF(NUM.LT.INC) INC=NUM	18810
	WRITE(6,30) (I,I=IB,INC)	18820
	WRITE(6,50)	18830
	K=-NO	18840
	DO 10 J=1,NUM	18850
	K=K+NO	18860
10	WRITE(6,40) J,(C(I+K),I=IB,INC)	18870
	WRITE(6,60)	18880
	IB=INC+1	18890
20	CONTINUE	18900
	RETURN	18910
25	NR=NO/3	18920
	IF((3*NR).NE.NO) NR=NR+1	18930
	DO 26 K=1,NR	18940
26	WRITE(6,80) (L,C(L),L=K,NO,NR)	18950
	RETURN	18960
30	FORMAT (1H0,8X,I3,9(9X,I3))	18970
40	FORMAT (1H ,I3,10(1X,E11.4))	18980
50	FORMAT (1H )	18990
60	FORMAT (1H0)	19000
80	FORMAT(1H ,3X,3(I3,5X,E11.4,5X))	19010
	END	19020
	SUBROUTINE UDU (NUMEQ,LU)	19030
C**	AN LDU DECOMPOSITION FOR BANDED SYMETRIC MATRICES	19040
	COMMON/LDUBLK/CPKC(490,180),CPKH(490, 90),SH(8, 490),	19050
	\$ QC( 490),QH( 490),IBC,IBH	19060
	REAL*8 CPKC,CPKH,SH,QC,QH,FAC	19070
	NEM1=NUMEQ-1	19080
	DO 450 I=1,NEM1	19090
	JEND=NUMEQ-I+1	19100
	IF(JEND.GT.IBH) JEND=IBH	19110
	DO 440 J=2,JEND	19120
	J1=I+J-1	19130
	IF(LU.EQ.0) GO TO 435	19140
	FAC=CPKH(I,J)/CPKH(I,1)	19150
429	K1=0	19160
	DO 430 K=J,JEND	19170
	K1=K1+1	19180
	CPKH(J1,K1)=CPKH(J1,K1)-CPKH(I,K)*FAC	19190
430	CONTINUE	19200
	CPKH(I,J)=FAC	19210
435	CONTINUE	19220
	QH(J1)=QH(J1)-QH(I)*CPKH(I,J)	19230
440	CONTINUE	19240

Table 2.--Computer-program listing--Continued

450	CONTINUE	19250
C	BACK SUBSTITUTION	19260
	QH(NUMEQ)=QH(NUMEQ)/CPKH(NUMEQ,1)	19270
	DO 470 I=1,NEM1	19280
	I1=NUMEQ-I	19290
	QH(I1)=QH(I1)/CPKH(I1,1)	19300
	JEND=NUMEQ-I1+1	19310
	IF(JEND.GT.IBH) JEND=IBH	19320
	DO 460 J=2,JEND	19330
	J1=I1+J-1	19340
	QH(I1)=QH(I1)-CPKH(I1,J)*QH(J1)	19350
460	CONTINUE	19360
470	CONTINUE	19370
	RETURN	19380
	END	19390
C		19400
	SUBROUTINE LDU (NUMEQ,LU)	19410
C**	AN LDU DECOMPOSITION FOR BANDED NONSYMMETRIC MATRICES	19420
	COMMON /LDUBLK/ CPKC(490,180),CPKH(490, 90),SH(8, 490),	19430
	\$ QC( 490),QH( 490),IBC,IBH	19440
	REAL*8 CPKC,CPKH,SH,QC,QH,FAC	19450
	NEM1=NUMEQ-1	19460
	IBD2=(IBC-1)/2	19470
	IDIAG=IBD2+1	19480
	DO 450 I=1,NEM1	19490
	JEND=NUMEQ-I	19500
	IF(JEND.GT.IBD2) JEND=IBD2	19510
	DO 440 J=1,JEND	19520
	JROW=I+J	19530
	JCOL=IBD2-J+1	19540
	IF(LU.EQ.0) GO TO 435	19550
	FAC=CPKC(JROW,JCOL)/CPKC(I,IDIAG)	19560
	CPKC(JROW,JCOL)=FAC	19570
	DO 430 K=1,JEND	19580
	K1=IDIAG+K	19590
	K2=JCOL+K	19600
	CPKC(JROW,K2)=CPKC(JROW,K2)-FAC*CPKC(I,K1)	19610
430	CONTINUE	19620
435	CONTINUE	19630
	QC(JROW)=QC(JROW)-CPKC(JROW,JCOL)*QC(I)	19640
440	CONTINUE	19650
450	CONTINUE	19660
C	BACK SUBSTITUTION	19670
	QC(NUMEQ)=QC(NUMEQ)/CPKC(NUMEQ,IDIAG)	19680
	JBGN=IDIAG+1	19690
	DO 470 I=1,NEM1	19700
	IROW=NUMEQ-I	19710
	JEND=IDIAG+I	19720
	IF(JEND.GT.IBC) JEND=IBC	19730
	JROW=IROW	19740
	DO 460 J=JBGN,JEND	19750
	JROW=JROW+1	19760

Table 2.--Computer-program listing--Continued

	QC(IROW)=QC(IROW)-CPKC(IROW,J)*QC(JROW)	19770
460	CONTINUE	19780
	QC(IROW)=QC(IROW)/CPKC(IROW,IDIAG)	19790
470	CONTINUE	19800
	RETURN	19810
	END	19820
C		19830
C**	SHAPE FACTOR SUBROUTINE	19840
C**	FOR THREE DIMENSIONAL CUBIC ELEMENTS	19850
C**	SF(FUNCTION,NODAL POINT,QUADRATURE POINT)	19860
	SUBROUTINE SHAFAC(THETAC)	19870
	DIMENSION SF(4,8,8),WF(4,8),WT(8)	19880
	DIMENSION QPT(8,3)	19890
	COMMON /SHABLK/ SF,WF,WT,NUMQPT	19900
	DOUBLE PRECISION SF,WF,WT,QPT,A,B,C	19910
	NUMQPT=8	19920
	WT(1)=1.0/8.0	19930
	WT(2)=1.0/8.0	19940
	WT(3)=WT(2)	19950
	WT(4)=WT(2)	19960
	WT(5)=WT(2)	19970
	WT(6)=WT(2)	19980
	WT(7)=WT(2)	19990
	WT(8)=WT(2)	20000
	QPT(1,1)=0.211324865405187	20010
	QPT(1,2)=0.211324865405187	20020
	QPT(1,3)=0.211324865405187	20030
	QPT(2,1)=0.788675134594813	20040
	QPT(2,2)=QPT(1,1)	20050
	QPT(2,3)=QPT(1,3)	20060
	QPT(3,1)=QPT(1,1)	20070
	QPT(3,2)=QPT(2,1)	20080
	QPT(3,3)=QPT(1,3)	20090
	QPT(4,1)=QPT(2,1)	20100
	QPT(4,2)=QPT(2,1)	20110
	QPT(4,3)=QPT(1,3)	20120
	QPT(5,1)=QPT(1,1)	20130
	QPT(5,2)=QPT(1,2)	20140
	QPT(5,3)=0.788675134594813	20150
	QPT(6,1)=QPT(2,1)	20160
	QPT(6,2)=QPT(2,2)	20170
	QPT(6,3)=QPT(5,3)	20180
	QPT(7,1)=QPT(3,1)	20190
	QPT(7,2)=QPT(3,2)	20200
	QPT(7,3)=QPT(5,3)	20210
	QPT(8,1)=QPT(4,1)	20220
	QPT(8,2)=QPT(4,2)	20230
	QPT(8,3)=QPT(5,3)	20240
	DO 200 I=1,NUMQPT	20250
	A=QPT(I,1)	20260
	B=QPT(I,2)	20270
	C=QPT(I,3)	20280

Table 2.--Computer-program listing--Continued

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SF(1,1,I)=(1.0-A)*(B-1.0)*(C-1.0)                20290
SF(1,2,I)=A*(B-1.0)*(C-1.0)                        20300
SF(1,3,I)=A*B*(1.0-C)                             20310
SF(1,4,I)=B*(A-1.0)*(C-1.0)                        20320
SF(1,5,I)=C*(A-1.0)*(B-1.0)                        20330
SF(1,6,I)=A*C*(1.0-B)                             20340
SF(1,7,I)=A*B*C                                    20350
SF(1,8,I)=B*C*(1.0-A)                             20360
SF(2,1,I)=(1.0-B)*(C-1.0)                          20370
SF(2,2,I)=(B-1.0)*(C-1.0)                          20380
SF(2,3,I)=B*(1.0-C)                                20390
SF(2,4,I)=B*(C-1.0)                                20400
SF(2,5,I)=C*(B-1.0)                                20410
SF(2,6,I)=C*(1.0-B)                                20420
SF(2,7,I)=B*C                                       20430
SF(2,8,I)=(-1.0)*B*C                               20440
SF(3,1,I)=(1.0-A)*(C-1.0)                          20450
SF(3,2,I)=A*(C-1.0)                                20460
SF(3,3,I)=(1.0-C)*A                                20470
SF(3,4,I)=(A-1.0)*(C-1.0)                          20480
SF(3,5,I)=C*(A-1.0)                                20490
SF(3,6,I)=(-1.0)*A*C                               20500
SF(3,7,I)=A*C                                       20510
SF(3,8,I)=C*(1.0-A)                                20520
SF(4,1,I)=(1.0-A)*(B-1.0)                          20530
SF(4,2,I)=A*(B-1.0)                                20540
SF(4,3,I)=(-1.0)*A*B                               20550
SF(4,4,I)=B*(A-1.0)                                20560
SF(4,5,I)=(A-1.0)*(B-1.0)                          20570
SF(4,6,I)=(-1.0)*A*(B-1.0)                        20580
SF(4,7,I)=A*B                                       20590
SF(4,8,I)=B*(1.0-A)                                20600
IF (THETAC.LT.0.0) GO TO 200                        20610
WF(1,I)=-108.0*A*B*C*(A-1.0)*(B-1.0)*(C-1.0)      20620
WF(2,I)=-108.0*B*C*(B-1.0)*(C-1.0)*(2.0*A-1.0)    20630
WF(3,I)=-108.0*A*C*(A-1.0)*(C-1.0)*(2.0*B-1.0)    20640
WF(4,I)=-108.0*A*B*(A-1.0)*(B-1.0)*(2.0*C-1.0)    20650
200  CONTINUE                                       20660
      RETURN                                       20670
      END                                         20680
C** TWO DIMENSIONAL SHAPE FACTOR ROUTINE           20690
      SUBROUTINE SURFAC                           20700
      COMMON/SURBLK/ AF(3,4),AWT,NQPTA            20710
      DOUBLE PRECISION AF,AWT                     20720
      NQPTA=1                                      20730
      AF(1,1)=0.25                                20740
      AF(1,2)=0.25                                20750
      AF(1,3)=0.25                                20760
      AF(1,4)=0.25                                20770
      AF(2,1)=-0.5                                20780
      AF(2,2)=+0.5                                20790
      AF(2,3)=+0.5                                20800

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Table 2.--*Computer-program listing*--Continued

AF(2,4)=-0.5	20810
AF(3,1)=-0.5	20820
AF(3,2)=-0.5	20830
AF(3,3)=+0.5	20840
AF(3,4)=+0.5	20850
AWT=1.0	20860
RETURN	20870
END	20880
SUBROUTINE LINFAC	20890
COMMON/LINBLK/RLF(2,2),RLWT,LINOD(4,2)	20900
LINOD(1,1)=1	20910
LINOD(1,2)=5	20920
LINOD(2,1)=2	20930
LINOD(2,2)=6	20940
LINOD(3,1)=3	20950
LINOD(3,2)=7	20960
LINOD(4,1)=4	20970
LINOD(4,2)=8	20980
RLWT=1.0	20990
RLF(1,1)=0.5	21000
RLF(2,1)=-1.0	21010
RLF(1,2)=0.5	21020
RLF(2,2)=1.0	21030
RETURN	21040
END	21050

**Table 3.--Instructions for preparing model data**

**Group 1: Title and problem setup**

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1-3	1-80	20A4	TITLE	Description of problem.
4	1-5	I5	NUMEL	Number of elements.
	6-10	I5	NUMNP	Number of nodes.
	11-15	I5	NUMAT	Number of material zones.
	16-20	I5	NTIME	Number of times steps in transport problem.
	21-25	I5	NTPER	Number of transport time steps per flow time step.
	26-30	I5	NTB	Number of times when boundary conditions change.
	31-35	I5	NTO	Number of times with observed water-level and concentration data.
	36-40	I5	NOBMAX	Number of observations for a single time step.
	41-45	I5	NTW	Number of times when point source-sinks change.
	46-50	I5	NTS	Number of times when distributed source-sinks change.
	51-55	I5	NPARH	Number of flow-calibration parameters.
	56-50	I5	NQPAR	Number of point and distributed source-sink parameters considered in calibration.
	66-70	I5	NBPAR	Number of specified head and concentration boundaries considered as calibration parameters.
5	1-5	I5	NPLAYR	Maximum number of nodes per horizontal layer.
	6-10	I5	NEPLAY	Maximum number of elements per horizontal layer.
6	1-5	I5	ITMAX	Maximum number of regression iterations for steady-state flow problem. If not using regression technique code 0.
	6-10	I5	IVELPR	Code 1 for printout of velocity and dispersion coefficients.
	11-15	I5	IFRAC	Code 1 to use dual-porosity option in transport problems (equation 5).
	16-20	I5	NTFRAC	Number of time steps used in equation 5 summation.
	21-25	I5	ITFRAC	Number of iterations used in equation 5 summation.
	26-30	I5	NLAYER	Maximum number of nodes in any vertical section.
	31-35	I5	IBEALE	Code 1 to compute coefficients for Beale's measure of nonlinearity.

Table 3.--*Instructions for preparing model data*--Continued

## Group 1: Title and problem setup--Continued

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
7	1-5	I5	NTPRT	Number of transport time steps between printouts.
	6-10	I5	IPO	Code 1 to print out sensitivity coefficients.
	11-15	I5	IPRX	Code 1 for additional printout of regression results.
	16-25	E10.3	AP	Acceleration parameter for regression analysis (1.0 if not used).
	26-35	E10.3	AMP	Marquardt parameter for regression analysis (0.0 if not used).
	36-45	E10.3	RP	Ridge parameter for regression analysis (0.0 if not used).
	46-55	E10.3	RPF	Bias parameter for regression analysis (0.0 if not used).
8	56-65	E10.3	EVH	Estimated error variance for flow problem (if not known code 0.0).
	1-10	E10.3	DTIME	Initial time step for transport problem.
	11-20	E10.3	TIMFAC	Time step expansion factor.
	21-30	E10.3	THETAH	$\theta$ in equation 2. Code 1.0 if simulating steady state.
	31-40	E10.3	THETAC	$\theta$ in equation 4. Code -1.0 if not modeling transport.
	41-50	E10.3	DFW	Molecular-dispersion coefficient within fractures.
	51-60	E10.3	CEQUI	Equilibrium solute concentration in equation 6.
9	1-10	E10.3	ALFLG	Code 0.0 if values for this card are to be computed by the program; otherwise code 1.0.
	11-20	E10.3	ALFAX	Upstream weighting factor in x direction.
	21-30	E10.3	ALFAY	Upstream weighting factor in y direction.
	31-40	E10.3	ALFAZ	Upstream weighting factor in z direction. Upstream weighting is useful in advection-dominated transport problems. For a complete discussion of upstream weights refer to Noorishad and Mehran (1982), and Huyakorn and Nilkuha (1979). For no upstream weighting code ALFAX = ALFAY = ALFAZ = 0.0.
10	1-70	7E10.3	ZSPACE (I), I=1,(NYLAYER-1)	Nodal spacing in vertical direction measured from bottom layer.

**Table 3.--Instructions for preparing model data--Continued**

**Group 2: Nodal locations**

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-10	I10	I	Node number.
	11-20	I10	N	Number of nodes directly overlying node I (include node I).
	21-30	E10.3	XORD	Distance in x-direction
	31-40	E10.3	YORD	Distance in y-direction.
	41-50	E10.3	ZORD	Distance in z-direction.

**Note:**

Only nodes in the bottom layer are coded. The location of nodes in other layers is calculated by the program. The x and y coordinates of overlying nodes are the same as those of the bottom node. Vertical spacing is determined by ZSPACE. A total of NPLAYR nodal-location cards are coded.

**Group 3: Nodal information for ground-water flow problem**

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-10	I10	N	Number of nodal flow cards to follow.
--	1-10	I10	I	Node number.
	11-20	I10	NPBC	Nodal-point boundary condition. Code 0 for no boundary condition, - 1 for specified head-boundary nodes, - 2 for specified concentration-boundary nodes, - 3 for specified head and concentration nodes, - 4 for boundary nodes without dispersive solute flux, or - 5 for boundary nodes with specified head and without dispersive solute flux.
	21-30	I10	NODOBS	Observation number. If node is not located along a well bore code 0.
	31-40	E10.3	WTNOD	Weight between 0.0 and 1.0 used to indicate partial penetration of an observation well. Code 1.0 if the well reaches the node, 0.0 if the well only reaches the overlying or underlying node.
	41-50	E10.3	H	Initial hydraulic head.
	51-60	E10.3	WELL	Initial point source-sink. If WELL is treated as a calibration parameter NPBC is used to represent the parameter number and must be coded greater than zero.
	61-70	E10.3	QS	Initial distributed source-sink.

Table 3.--*Instructions for preparing model data*--Continued

Group 4: Nodal information for transport problem

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-10	I10	N	Number of cards to follow. Do not code this card if only solving a flow problem.
--	1-10	I10	I	Node number.
	11-20	E10.3	C	Initial concentration.
	21-30	E10.3	WELLC	Initial concentration of point source.
	31-40	E10.3	CS	Initial concentration of distributed source.

Group 5: Nodal configuration of each element

Only the basal NEPLAY elements are coded.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-5	I5	I	Element number.
	6-10	I5	N	Number of elements directly overlying element I (include basal element).
	11-30	4I5	NP	Nodes associated with element I. Only the four basal nodes of element I are coded in counter-clockwise direction, as viewed from above.

Group 6: Aquifer zonal configuration

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-10	I5	N	Numbers of cards to follow.
--	1-10	I10	I	Element number.
	11-20	I10	MAT	Material-zone number associated with element I. Code as the negative of the zone number if a distributed source is defined over the element. If no card is coded for an element, MAT is set equal to the element layer, numbered from the base of the problem.
	21-30	I10	IBPRM	Calibration-parameter number for distributed source-sinks.

Table 3.--*Instructions for preparing model data*--Continued

Group 7: Aquifer parameters for ground-water flow

One card must be coded for each zone.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-10	I10	I	Material-zone number.
	11-20	E10.3	RX	Hydraulic conductivity in the x-direction.
	21-30	E10.3	RY	Hydraulic conductivity in the y-direction.
	31-40	E10.3	RZ	Hydraulic conductivity in the z-direction.
	41-50	E10.3	RC	Specific storage.

Group 8: Aquifer parameters for solute transport

One card must be coded for each zone, ordered by the zone number. If not solving a transport problem this group is not coded.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-10	E10.3	DL	Longitudinal dispersivity.
	11-20	E10.3	DT	Transverse dispersivity.
	21-30	E10.3	DM	Molecular-diffusion coefficient within the media blocks. Only used if simulating dual-porosity media.
	31-40	E10.3	RPR	Porosity. If simulating dual-porosity media, porosity of fractures.
	41-50	E10.3	FP	Fracture width.
	51-60	E10.3	FQ	Fracture spacing or block width.
	61-70	E10.3	RMT	Mass-transfer coefficient $K^*A$ in equation 6. Code 0.0 if not used.
	71-80	E10.3	RPBR	Block porosity in dual-porosity media.

Group 9: Initial solute concentration of water within the media blocks

Code only if simulating dual-porosity media. One value is coded per material zone, ordered by zone number.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-80	8E10.3	CBINIT	Initial concentration of water within the media blocks.

Group 10: Calibration-parameter numbers for ground-water flow

If NPARH = 0 do not code these cards.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-5	I5	I	Material-zone number.
	6-10	I5	IPRMH(1,I)	Parameter number for RX. Code 0 if RX is not a calibration parameter.

Table 3.--*Instructions for preparing model data*--Continued

Group 10: Calibration-parameter numbers for ground-water flow--Continued

If NPARH = 0 do not code these cards.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
	11-15	I5	IPRMH(2,I)	Parameter number for RY.
	16-20	I5	IPRMH(3,I)	Parameter number for RZ.

Group 11: Standard errors of calibration parameters for  
ground-water flow

One card per calibration parameter is coded. If NPARH is coded as 0 do not code these cards.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-10	I10	I	Parameter number.
	11-20	E10.3	RKH	Normalized standard error of estimate for parameter I. Code 0.0 if no prior information exists for the parameter.

Group 12: Information for source-sink calibration parameters in  
the ground-water flow problem

Code one card per source-sink parameter.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-10	I10	I	Source-sink parameter number.
	11-20	E10.3	QF	Multiplication factor for source-sink (either WELL or QS).
	21-30	E10.3	RKH	Normalized standard error for parameter.

Group 13: Boundary-parameter number associated with each node

Coded 15 values per card ordered by node number. If NBPAP is coded as 0 do not code these cards.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
-- Boundary	1-75	15I5	NBP	Boundary parameter number for each node.

Table 3.--*Instructions for preparing model data*--ContinuedGroup 14: Standard errors of boundary parameters for ground-water flow

Code NBPARG cards. If NBPARG is coded as 0 do not code these cards.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
--	1-10	I10 E10.3	I RKH	Boundary-parameter number. Normalized standard error of head-boundary parameter I.

Group 15: Transient boundary condition cards

NTB sets of these cards are needed.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-5	I5	NXTB	Transport time step when boundary conditions change.
	6-10	I5	I2	Number of cards that follow.
--	1-5	I5	J	Node number.
	6-10	I5	NPBC	New nodal-point boundary condition.
	11-20	E10.3	H	New head.
	21-30	E10.3	C	New concentration.

Group 16: Observation cards

NTO sets of these cards are needed.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-10	I10	NXTO	Transport time step corresponding to the following observations.
--	1-10	I10	K	Observation number corresponding to NODOBS. NOBMAX nodal observation cards must be coded for each set.
	11-20	E10.3	HO	Observed water level.
	21-30	E10.3	WH	Weighting value for HO.
	31-40	E10.3	CO	Observed concentration.
	41-50	E10.3	WC	Weighting value for CO.

Group 17: Transient point source-sink cards

NTW sets of these cards are needed.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-10	I10	NXTW	Transport time step when source-sink changes.
	11-20	I10	I2	Number of cards that follow.
--	1-10	I10	K	Node number.
	11-20	E10.0	WELL	New point source-sink.
	21-30	E10.0	WELLC	New point-source concentration.



Table 3.--*Instructions for preparing model data*--Continued

Group 18: Transient distributed source-sink cards

NTS sets of these cards are needed.

<u>Card</u>	<u>Column</u>	<u>Format</u>	<u>Variable</u>	<u>Description</u>
1	1-10	I10	NXTS	Transport time step when source-sink changes.
	11-20	I10	I2	Number of cards that follow.
--	1-10	I10	K	Node number.
	11-20	E10.3	QS	New distributed source-sink.
	21-30	E10.3	CS	New distributed source concentration.

---

Table 4.--Input data for sample simulation

300	484	3	05	05	0	0	0	0	0	0	0	0
121	100											
0	1	1	8	10	4	0						
1	0	0		1.0		0.0		0.0		0.0		0.0
100.0			1.0		1.0		+1.0		0.0		0.0	
+1.0			0.0		0.0		0.0					
200.0			400.0		500.0							
1			4		0.00		0.00		0.0			
5			4		0.00		500.00		0.0			
9			4		0.00		1000.00		0.0			
13			4		0.00		1500.00		0.0			
17			4		0.00		2000.00		0.0			
21			4		0.00		2500.00		0.0			
25			4		0.00		3000.00		0.0			
29			4		0.00		3500.00		0.0			
33			4		0.00		4000.00		0.0			
37			4		0.00		4500.00		0.0			
41			4		0.00		5000.00		0.0			
45			4		500.00		0.00		0.0			
49			4		500.00		500.00		0.0			
53			4		500.00		1000.00		0.0			
57			4		500.00		1500.00		0.0			
61			4		500.00		2000.00		0.0			
65			4		500.00		2500.00		0.0			
69			4		500.00		3000.00		0.0			
73			4		500.00		3500.00		0.0			
77			4		500.00		4000.00		0.0			
81			4		500.00		4500.00		0.0			
85			4		500.00		5000.00		0.0			
89			4		1000.00		0.00		0.0			
93			4		1000.00		500.00		0.0			
97			4		1000.00		1000.00		0.0			
101			4		1000.00		1500.00		0.0			
105			4		1000.00		2000.00		0.0			
109			4		1000.00		2500.00		0.0			
113			4		1000.00		3000.00		0.0			
117			4		1000.00		3500.00		0.0			
121			4		1000.00		4000.00		0.0			
125			4		1000.00		4500.00		0.0			
129			4		1000.00		5000.00		0.0			
133			4		1500.00		0.00		0.0			
137			4		1500.00		500.00		0.0			
141			4		1500.00		1000.00		0.0			
145			4		1500.00		1500.00		0.0			
149			4		1500.00		2000.00		0.0			
153			4		1500.00		2500.00		0.0			
157			4		1500.00		3000.00		0.0			
161			4		1500.00		3500.00		0.0			
165			4		1500.00		4000.00		0.0			
169			4		1500.00		4500.00		0.0			
173			4		1500.00		5000.00		0.0			
177			4		2000.00		0.00		0.0			

Table 4.--*Input data for sample simulation*--Continued

181	4	2000.00	500.00	0.0
185	4	2000.00	1000.00	0.0
189	4	2000.00	1500.00	0.0
193	4	2000.00	2000.00	0.0
197	4	2000.00	2500.00	0.0
201	4	2000.00	3000.00	0.0
205	4	2000.00	3500.00	0.0
209	4	2000.00	4000.00	0.0
213	4	2000.00	4500.00	0.0
217	4	2000.00	5000.00	0.0
221	4	2500.00	0.00	0.0
225	4	2500.00	500.00	0.0
229	4	2500.00	1000.00	0.0
233	4	2500.00	1500.00	0.0
237	4	2500.00	2000.00	0.0
241	4	2500.00	2500.00	0.0
245	4	2500.00	3000.00	0.0
249	4	2500.00	3500.00	0.0
253	4	2500.00	4000.00	0.0
257	4	2500.00	4500.00	0.0
261	4	2500.00	5000.00	0.0
265	4	3000.00	0.00	0.0
269	4	3000.00	500.00	0.0
273	4	3000.00	1000.00	0.0
277	4	3000.00	1500.00	0.0
281	4	3000.00	2000.00	0.0
285	4	3000.00	2500.00	0.0
289	4	3000.00	3000.00	0.0
293	4	3000.00	3500.00	0.0
297	4	3000.00	4000.00	0.0
301	4	3000.00	4500.00	0.0
305	4	3000.00	5000.00	0.0
309	4	3500.00	0.00	0.0
313	4	3500.00	500.00	0.0
317	4	3500.00	1000.00	0.0
321	4	3500.00	1500.00	0.0
325	4	3500.00	2000.00	0.0
329	4	3500.00	2500.00	0.0
333	4	3500.00	3000.00	0.0
337	4	3500.00	3500.00	0.0
341	4	3500.00	4000.00	0.0
345	4	3500.00	4500.00	0.0
349	4	3500.00	5000.00	0.0
353	4	4000.00	0.00	0.0
357	4	4000.00	500.00	0.0
361	4	4000.00	1000.00	0.0
365	4	4000.00	1500.00	0.0
369	4	4000.00	2000.00	0.0
373	4	4000.00	2500.00	0.0
377	4	4000.00	3000.00	0.0
381	4	4000.00	3500.00	0.0
385	4	4000.00	4000.00	0.0

Table 4.--Input data for sample simulation--Continued

389	4	4000.00	4500.00	0.0		
393	4	4000.00	5000.00	0.0		
397	4	4500.00	0.00	0.0		
401	4	4500.00	500.00	0.0		
405	4	4500.00	1000.00	0.0		
409	4	4500.00	1500.00	0.0		
413	4	4500.00	2000.00	0.0		
417	4	4500.00	2500.00	0.0		
421	4	4500.00	3000.00	0.0		
425	4	4500.00	3500.00	0.0		
429	4	4500.00	4000.00	0.0		
433	4	4500.00	4500.00	0.0		
437	4	4500.00	5000.00	0.0		
441	4	5000.00	0.00	0.0		
445	4	5000.00	500.00	0.0		
449	4	5000.00	1000.00	0.0		
453	4	5000.00	1500.00	0.0		
457	4	5000.00	2000.00	0.0		
461	4	5000.00	2500.00	0.0		
465	4	5000.00	3000.00	0.0		
469	4	5000.00	3500.00	0.0		
473	4	5000.00	4000.00	0.0		
477	4	5000.00	4500.00	0.0		
481	4	5000.00	5000.00	0.0		
59						
1	0	0	0.0	0.0	0.0	6.0E-01
2	0	0	0.0	0.0	0.0	2.0E-01
3	0	0	0.0	0.0	0.0	2.0E-01
4	0	0	0.0	0.0	0.0	6.0E-01
5	0	0	0.0	0.0	0.0	6.0E-01
6	0	0	0.0	0.0	0.0	2.0E-01
7	0	0	0.0	0.0	0.0	2.0E-01
8	0	0	0.0	0.0	0.0	6.0E-01
9	0	0	0.0	0.0	0.0	6.0E-01
10	0	0	0.0	0.0	0.0	2.0E-01
11	0	0	0.0	0.0	0.0	2.0E-01
12	0	0	0.0	0.0	0.0	6.0E-01
13	0	0	0.0	0.0	0.0	6.0E-01
14	0	0	0.0	0.0	0.0	2.0E-01
15	0	0	0.0	0.0	0.0	2.0E-01
16	0	0	0.0	0.0	0.0	6.0E-01
17	0	0	0.0	0.0	0.0	6.0E-01
18	0	0	0.0	0.0	0.0	2.0E-01
19	0	0	0.0	0.0	0.0	2.0E-01
20	0	0	0.0	0.0	0.0	6.0E-01
21	0	0	0.0	0.0	0.0	6.0E-01
22	0	0	0.0	0.0	0.0	2.0E-01
23	0	0	0.0	0.0	0.0	2.0E-01
24	0	0	0.0	0.0	0.0	6.0E-01
25	0	0	0.0	0.0	0.0	6.0E-01
26	0	0	0.0	0.0	0.0	2.0E-01
27	0	0	0.0	0.0	0.0	2.0E-01

Table 4.--Input data for sample simulation--Continued

28	0	0	0.0	0.0	0.0	6.0E-01
29	0	0	0.0	0.0	0.0	6.0E-01
30	0	0	0.0	0.0	0.0	2.0E-01
31	0	0	0.0	0.0	0.0	2.0E-01
32	0	0	0.0	0.0	0.0	6.0E-01
33	0	0	0.0	0.0	0.0	6.0E-01
34	0	0	0.0	0.0	0.0	2.0E-01
35	0	0	0.0	0.0	0.0	2.0E-01
36	0	0	0.0	0.0	0.0	6.0E-01
37	0	0	0.0	0.0	0.0	6.0E-01
38	0	0	0.0	0.0	0.0	2.0E-01
39	0	0	0.0	0.0	0.0	2.0E-01
40	0	0	0.0	0.0	0.0	6.0E-01
41	0	0	0.0	0.0	0.0	6.0E-01
42	0	0	0.0	0.0	0.0	2.0E-01
43	0	0	0.0	0.0	0.0	2.0E-01
44	0	0	0.0	0.0	0.0	6.0E-01
444	-5	0	0.0	500.0	0.0	0.0
448	-5	0	0.0	500.0	0.0	0.0
452	-5	0	0.0	500.0	0.0	0.0
456	-5	0	0.0	500.0	0.0	0.0
460	-5	0	0.0	500.0	0.0	0.0
464	-5	0	0.0	500.0	0.0	0.0
468	-5	0	0.0	500.0	0.0	0.0
472	-5	0	0.0	500.0	0.0	0.0
476	-5	0	0.0	500.0	0.0	0.0
480	-5	0	0.0	500.0	0.0	0.0
484	-5	0	0.0	500.0	0.0	0.0
161	0	0	0.0	0.0	-50000.0	0.0
162	0	0	0.0	0.0	-50000.0	0.0
323	0	0	0.0	0.0	-25000.0	0.0
324	0	0	0.0	0.0	-25000.0	0.0
4						
9	0.0	0.0	100.0			
10	0.0	0.0	100.0			
13	0.0	0.0	100.0			
14	0.0	0.0	100.0			
1	3	1	45	49	5	
4	3	5	49	53	9	
7	3	9	53	57	13	
10	3	13	57	61	17	
13	3	17	61	65	21	
16	3	21	65	69	25	
19	3	25	69	73	29	
22	3	29	73	77	33	
25	3	33	77	81	37	
28	3	37	81	85	41	
31	3	45	89	93	49	
34	3	49	93	97	53	
37	3	53	97	101	57	
40	3	57	101	105	61	
43	3	61	105	109	65	

Table 4.--*Input data for sample simulation*--Continued

46	3	65	109	113	69
49	3	69	113	117	73
52	3	73	117	121	77
55	3	77	121	125	81
58	3	81	125	129	85
61	3	89	133	137	93
64	3	93	137	141	97
67	3	97	141	145	101
70	3	101	145	149	105
73	3	105	149	153	109
76	3	109	153	157	113
79	3	113	157	161	117
82	3	117	161	165	121
85	3	121	165	169	125
88	3	125	169	173	129
91	3	133	177	181	137
94	3	137	181	185	141
97	3	141	185	189	145
100	3	145	189	193	149
103	3	149	193	197	153
106	3	153	197	201	157
109	3	157	201	205	161
112	3	161	205	209	165
115	3	165	209	213	169
118	3	169	213	217	173
121	3	177	221	225	181
124	3	181	225	229	185
127	3	185	229	233	189
130	3	189	233	237	193
133	3	193	237	241	197
136	3	197	241	245	201
139	3	201	245	249	205
142	3	205	249	253	209
145	3	209	253	257	213
148	3	213	257	261	217
151	3	221	265	269	225
154	3	225	269	273	229
157	3	229	273	277	233
160	3	233	277	281	237
163	3	237	281	285	241
166	3	241	285	289	245
169	3	245	289	293	249
172	3	249	293	297	253
175	3	253	297	301	257
178	3	257	301	305	261
181	3	265	309	313	269
184	3	269	313	317	273
187	3	273	317	321	277
190	3	277	321	325	281
193	3	281	325	329	285
196	3	285	329	333	289
199	3	289	333	337	293

Table 4.--*Input data for sample simulation*--Continued

202	3	293	337	341	297
205	3	297	341	345	301
208	3	301	345	349	305
211	3	309	353	357	313
214	3	313	357	361	317
217	3	317	361	365	321
220	3	321	365	369	325
223	3	325	369	373	329
226	3	329	373	377	333
229	3	333	377	381	337
232	3	337	381	385	341
235	3	341	385	389	345
238	3	345	389	393	349
241	3	353	397	401	357
244	3	357	401	405	361
247	3	361	405	409	365
250	3	365	409	413	369
253	3	369	413	417	373
256	3	373	417	421	377
259	3	377	421	425	381
262	3	381	425	429	385
265	3	385	429	433	389
268	3	389	433	437	393
271	3	397	441	445	401
274	3	401	445	449	405
277	3	405	449	453	409
280	3	409	453	457	413
283	3	413	457	461	417
286	3	417	461	465	421
289	3	421	465	469	425
292	3	425	469	473	429
295	3	429	473	477	433
298	3	433	477	481	437
15					
	1		-1		0
	2		-2		0
	3		-3		0
	4		-1		0
	5		-2		0
	6		-3		0
	7		-1		0
	8		-2		0
	9		-3		0
	10		-1		0
	11		-2		0
	12		-3		0
	13		-1		0
	14		-2		0
	15		-3		0
	1	25.0		25.0	5.0
	2	10.0		2.0	10.0
	3	40.0		40.0	0.0

Table 4.--*Input data for sample simulation*--Continued

10.0	3.0	1.0E-04	0.1	0.2	1.8	0.0	0.01
1.0	0.3	1.0E-04	0.05	0.01	0.19	0.0	0.01
100.0	30.0	0.0	0.3	0.0	0.0	0.0	0.0
0.0	0.0	0.0					



Table 5.--Selected output for sample simulation

OELEMENT	VX	VY	VZ	DISPX	DISPY	DISPZ
1	0.36793E+00	0.22678E-01	0.41290E-02	0.36764E+01	0.11157E+01	0.11063E+01
2	0.14245E+00	0.18094E-02	0.37103E-01	0.14065E+00	0.44179E-01	0.50709E-01
3	0.50986E+00	0.36154E-01	0.34060E-01	0.50890E+02	0.15547E+02	0.15527E+02
4	0.36669E+00	0.70220E-01	0.40384E-02	0.36410E+01	0.12126E+01	0.11204E+01
5	0.14196E+00	0.55819E-02	0.37090E-01	0.14013E+00	0.44199E-01	0.50609E-01
6	0.50792E+00	0.11171E+00	0.34075E-01	0.50285E+02	0.17311E+02	0.15791E+02
7	0.36299E+00	0.12519E+00	0.38113E-02	0.35540E+01	0.14377E+01	0.11522E+01
8	0.14050E+00	0.10006E-01	0.36993E-01	0.13858E+00	0.44172E-01	0.50269E-01
9	0.50217E+00	0.19927E+00	0.34304E-01	0.48848E+02	0.21375E+02	0.16393E+02
10	0.35497E+00	0.19852E+00	0.30204E-02	0.33888E+01	0.18985E+01	0.12203E+01
11	0.13749E+00	0.15444E-01	0.36990E-01	0.13536E+00	0.44131E-01	0.49653E-01
12	0.48985E+00	0.31212E+00	0.34831E-01	0.46323E+02	0.29176E+02	0.17602E+02
13	0.31312E+00	0.36325E+00	0.59700E-03	0.28698E+01	0.33647E+01	0.14387E+01
14	0.12257E+00	0.29051E-01	0.34528E-01	0.11970E+00	0.43707E-01	0.45574E-01
15	0.43271E+00	0.55446E+00	0.36048E-01	0.39739E+02	0.51684E+02	0.21257E+02
16	0.16265E+00	0.36199E+00	0.32516E-02	0.16572E+01	0.35019E+01	0.11908E+01
17	0.62525E-01	0.28962E-01	0.14318E-01	0.59997E-01	0.29457E-01	0.23153E-01
18	0.22703E+00	0.55271E+00	0.11141E-01	0.23966E+02	0.53711E+02	0.17943E+02
19	0.41392E-01	0.19312E+00	0.46153E-02	0.65340E+00	0.19142E+01	0.59345E+00
20	0.15033E-01	0.15020E-01	-0.11152E-02	0.13818E-01	0.13805E-01	0.64249E-02
21	0.63782E-01	0.30366E+00	-0.68293E-02	0.10228E+02	0.30108E+02	0.93213E+01
22	0.31771E-01	0.11402E+00	0.16639E-02	0.41482E+00	0.11239E+01	0.35529E+00
23	0.12650E-01	0.91000E-02	0.58970E-03	0.11862E-01	0.83955E-02	0.46940E-02
24	0.48325E-01	0.18118E+00	-0.96020E-03	0.64973E+01	0.17880E+02	0.56258E+01
25	0.21817E-01	0.58068E-01	0.14778E-02	0.23985E+00	0.56655E+00	0.18639E+00
26	0.86806E-02	0.45777E-02	0.10892E-03	0.83188E-02	0.44389E-02	0.29451E-02
27	0.35092E-01	0.91639E-01	-0.15176E-02	0.38226E+01	0.89340E+01	0.29459E+01
28	0.16481E-01	0.17336E-01	0.10127E-02	0.15124E+00	0.15969E+00	0.72122E-01
29	0.63071E-02	0.13643E-02	0.13207E-03	0.62506E-02	0.21382E-02	0.19382E-02
30	0.24624E-01	0.27225E-01	-0.92102E-03	0.22575E+01	0.25145E+01	0.11032E+01
31	0.31226E+00	0.21581E-01	0.21722E-02	0.31196E+01	0.94945E+00	0.93914E+00
32	0.12219E+00	0.17202E-02	0.82262E-03	0.12218E+00	0.36678E-01	0.36664E-01
33	0.48846E+00	0.34404E-01	-0.25443E-02	0.48798E+02	0.14860E+02	0.14691E+02
34	0.30894E+00	0.66361E-01	0.21089E-02	0.30623E+01	0.10455E+01	0.94808E+00
35	0.12088E+00	0.52933E-02	0.80482E-03	0.12084E+00	0.36463E-01	0.36305E-01
36	0.48326E+00	0.10579E+00	-0.24507E-02	0.47886E+02	0.16425E+02	0.14842E+02
37	0.29959E+00	0.11644E+00	0.19215E-02	0.29189E+01	0.12596E+01	0.96436E+00
38	0.11722E+00	0.92597E-02	0.76553E-03	0.11708E+00	0.35787E-01	0.35280E-01
39	0.46864E+00	0.18529E+00	-0.23205E-02	0.45625E+02	0.19887E+02	0.15119E+02
40	0.27795E+00	0.17957E+00	0.15094E-02	0.26270E+01	0.16748E+01	0.99278E+00
41	0.10889E+00	0.14315E-01	0.61683E-03	0.10852E+00	0.34255E-01	0.32951E-01
42	0.43542E+00	0.28514E+00	-0.19919E-02	0.41113E+02	0.26549E+02	0.15615E+02
43	0.23097E+00	0.23083E+00	0.12852E-02	0.21233E+01	0.21218E+01	0.97967E+00
44	0.90602E-01	0.18471E-01	-0.12718E-04	0.89883E-01	0.30323E-01	0.27740E-01
45	0.36330E+00	0.36977E+00	-0.26759E-02	0.33374E+02	0.34015E+02	0.15553E+02
46	0.17144E+00	0.23511E+00	0.12712E-02	0.15800E+01	0.22027E+01	0.87297E+00
47	0.66391E-01	0.18782E-01	0.48994E-05	0.65417E-01	0.24278E-01	0.20699E-01
48	0.26362E+00	0.37462E+00	-0.17797E-02	0.24362E+02	0.35188E+02	0.13743E+02
49	0.13266E+00	0.17537E+00	0.13104E-02	0.12199E+01	0.16387E+01	0.65975E+00
50	0.49120E-01	0.14109E-01	-0.53863E-03	0.48378E-01	0.18059E-01	0.15337E-01

Table 5.--Selected output for sample simulation--Continued

TIME STEP	1	TOTAL TIME OF SIMULATION			100.00		
	NODE	XORD	YORD	ZORD	H	C	
1	0.00000E+00	0.00000E+00	0.00000E+00	0.54688E+03	0.13656E+01		
2	0.00000E+00	0.00000E+00	0.20000E+03	0.54670E+03	0.13071E+01		
3	0.00000E+00	0.00000E+00	0.40000E+03	0.54577E+03	0.19450E+00		
4	0.00000E+00	0.00000E+00	0.50000E+03	0.54566E+03	0.22148E+00		
5	0.00000E+00	0.50000E+03	0.00000E+00	0.54642E+03	0.21257E+02		
6	0.00000E+00	0.50000E+03	0.20000E+03	0.54625E+03	0.17615E+02		
7	0.00000E+00	0.50000E+03	0.40000E+03	0.54531E+03	0.25413E+01		
8	0.00000E+00	0.50000E+03	0.50000E+03	0.54520E+03	0.73561E+00		
9	0.00000E+00	0.10000E+04	0.00000E+00	0.54500E+03	0.61517E+02		
10	0.00000E+00	0.10000E+04	0.20000E+03	0.54484E+03	0.50354E+02		
11	0.00000E+00	0.10000E+04	0.40000E+03	0.54390E+03	0.71874E+01		
12	0.00000E+00	0.10000E+04	0.50000E+03	0.54379E+03	0.17069E+01		
13	0.00000E+00	0.15000E+04	0.00000E+00	0.54247E+03	0.64242E+02		
14	0.00000E+00	0.15000E+04	0.20000E+03	0.54230E+03	0.52123E+02		
15	0.00000E+00	0.15000E+04	0.40000E+03	0.54138E+03	0.72752E+01		
16	0.00000E+00	0.15000E+04	0.50000E+03	0.54127E+03	0.18243E+01		
17	0.00000E+00	0.20000E+04	0.00000E+00	0.53841E+03	0.26126E+02		
18	0.00000E+00	0.20000E+04	0.20000E+03	0.53843E+03	0.20814E+02		
19	0.00000E+00	0.20000E+04	0.40000E+03	0.53744E+03	0.27289E+01		
20	0.00000E+00	0.20000E+04	0.50000E+03	0.53731E+03	0.95652E+00		
21	0.00000E+00	0.25000E+04	0.00000E+00	0.53054E+03	0.53993E+01		
22	0.00000E+00	0.25000E+04	0.20000E+03	0.53042E+03	0.37372E+01		
23	0.00000E+00	0.25000E+04	0.40000E+03	0.52995E+03	0.31057E+00		
24	0.00000E+00	0.25000E+04	0.50000E+03	0.52989E+03	0.34697E+00		
25	0.00000E+00	0.30000E+04	0.00000E+00	0.52268E+03	0.16408E+01		
26	0.00000E+00	0.30000E+04	0.20000E+03	0.52244E+03	0.51514E+00		
27	0.00000E+00	0.30000E+04	0.40000E+03	0.52249E+03	-0.22495E-01		
28	0.00000E+00	0.30000E+04	0.50000E+03	0.52251E+03	0.76086E-02		
29	0.00000E+00	0.35000E+04	0.00000E+00	0.51874E+03	0.95041E-01		
30	0.00000E+00	0.35000E+04	0.20000E+03	0.51868E+03	0.17359E-03		
31	0.00000E+00	0.35000E+04	0.40000E+03	0.51865E+03	-0.29428E-02		
32	0.00000E+00	0.35000E+04	0.50000E+03	0.51865E+03	-0.36870E-02		
33	0.00000E+00	0.40000E+04	0.00000E+00	0.51642E+03	-0.58869E-02		
34	0.00000E+00	0.40000E+04	0.20000E+03	0.51635E+03	0.37361E-03		
35	0.00000E+00	0.40000E+04	0.40000E+03	0.51635E+03	0.25604E-03		
36	0.00000E+00	0.40000E+04	0.50000E+03	0.51636E+03	0.53348E-03		
37	0.00000E+00	0.45000E+04	0.00000E+00	0.51523E+03	0.94322E-03		
38	0.00000E+00	0.45000E+04	0.20000E+03	0.51518E+03	-0.45226E-04		
39	0.00000E+00	0.45000E+04	0.40000E+03	0.51517E+03	-0.31195E-04		
40	0.00000E+00	0.45000E+04	0.50000E+03	0.51518E+03	-0.12459E-03		
41	0.00000E+00	0.50000E+04	0.00000E+00	0.51486E+03	-0.37950E-03		
42	0.00000E+00	0.50000E+04	0.20000E+03	0.51482E+03	0.85539E-05		
43	0.00000E+00	0.50000E+04	0.40000E+03	0.51482E+03	0.95106E-05		
44	0.00000E+00	0.50000E+04	0.50000E+03	0.51482E+03	0.59893E-04		
45	0.50000E+03	0.00000E+00	0.00000E+00	0.53951E+03	0.12931E+01		
46	0.50000E+03	0.00000E+00	0.20000E+03	0.53938E+03	0.10619E+01		
47	0.50000E+03	0.00000E+00	0.40000E+03	0.53936E+03	0.13798E+00		
48	0.50000E+03	0.00000E+00	0.50000E+03	0.53937E+03	0.12620E+00		
49	0.50000E+03	0.50000E+03	0.00000E+00	0.53906E+03	0.46206E+01		
50	0.50000E+03	0.50000E+03	0.20000E+03	0.53894E+03	0.36959E+01		

Table 5.--*Selected output for sample simulation*--Continued

51	0.50000E+03	0.50000E+03	0.40000E+03	0.53892E+03	0.44603E+00
52	0.50000E+03	0.50000E+03	0.50000E+03	0.53893E+03	0.42936E+00
53	0.50000E+03	0.10000E+04	0.00000E+00	0.53770E+03	0.11358E+02
54	0.50000E+03	0.10000E+04	0.20000E+03	0.53759E+03	0.88935E+01
55	0.50000E+03	0.10000E+04	0.40000E+03	0.53757E+03	0.10297E+01
56	0.50000E+03	0.10000E+04	0.50000E+03	0.53758E+03	0.10160E+01
57	0.50000E+03	0.15000E+04	0.00000E+00	0.53530E+03	0.13174E+02
58	0.50000E+03	0.15000E+04	0.20000E+03	0.53521E+03	0.10009E+02
59	0.50000E+03	0.15000E+04	0.40000E+03	0.53519E+03	0.11104E+01
60	0.50000E+03	0.15000E+04	0.50000E+03	0.53520E+03	0.11133E+01
61	0.50000E+03	0.20000E+04	0.00000E+00	0.53153E+03	0.77139E+01
62	0.50000E+03	0.20000E+04	0.20000E+03	0.53145E+03	0.56143E+01
63	0.50000E+03	0.20000E+04	0.40000E+03	0.53146E+03	0.59393E+00
64	0.50000E+03	0.20000E+04	0.50000E+03	0.53147E+03	0.59816E+00
65	0.50000E+03	0.25000E+04	0.00000E+00	0.52663E+03	0.34570E+01
66	0.50000E+03	0.25000E+04	0.20000E+03	0.52656E+03	0.21757E+01
67	0.50000E+03	0.25000E+04	0.40000E+03	0.52656E+03	0.17815E+00
68	0.50000E+03	0.25000E+04	0.50000E+03	0.52657E+03	0.19774E+00
69	0.50000E+03	0.30000E+04	0.00000E+00	0.52179E+03	0.70828E+00
70	0.50000E+03	0.30000E+04	0.20000E+03	0.52168E+03	0.17924E+00
71	0.50000E+03	0.30000E+04	0.40000E+03	0.52165E+03	-0.19552E-01
72	0.50000E+03	0.30000E+04	0.50000E+03	0.52165E+03	0.32101E-02
73	0.50000E+03	0.35000E+04	0.00000E+00	0.51806E+03	0.39377E-01
74	0.50000E+03	0.35000E+04	0.20000E+03	0.51799E+03	0.30788E-02
75	0.50000E+03	0.35000E+04	0.40000E+03	0.51801E+03	0.22454E-03
76	0.50000E+03	0.35000E+04	0.50000E+03	0.51802E+03	-0.28177E-02
77	0.50000E+03	0.40000E+04	0.00000E+00	0.51596E+03	-0.34128E-02
78	0.50000E+03	0.40000E+04	0.20000E+03	0.51589E+03	-0.11398E-03
79	0.50000E+03	0.40000E+04	0.40000E+03	0.51588E+03	-0.19108E-03
80	0.50000E+03	0.40000E+04	0.50000E+03	0.51589E+03	0.55248E-03
81	0.50000E+03	0.45000E+04	0.00000E+00	0.51488E+03	0.53509E-03
82	0.50000E+03	0.45000E+04	0.20000E+03	0.51485E+03	0.11777E-04
83	0.50000E+03	0.45000E+04	0.40000E+03	0.51486E+03	0.46362E-04
84	0.50000E+03	0.45000E+04	0.50000E+03	0.51486E+03	-0.13666E-03
85	0.50000E+03	0.50000E+04	0.00000E+00	0.51460E+03	-0.21623E-03
86	0.50000E+03	0.50000E+04	0.20000E+03	0.51456E+03	-0.55297E-05
87	0.50000E+03	0.50000E+04	0.40000E+03	0.51456E+03	-0.16452E-04
88	0.50000E+03	0.50000E+04	0.50000E+03	0.51456E+03	0.60816E-04
89	0.10000E+04	0.00000E+00	0.00000E+00	0.53322E+03	0.29068E+00
90	0.10000E+04	0.00000E+00	0.20000E+03	0.53326E+03	0.17762E+00
91	0.10000E+04	0.00000E+00	0.40000E+03	0.53325E+03	0.91995E-02
92	0.10000E+04	0.00000E+00	0.50000E+03	0.53325E+03	0.14266E-01
93	0.10000E+04	0.50000E+03	0.00000E+00	0.53282E+03	0.78351E+00
94	0.10000E+04	0.50000E+03	0.20000E+03	0.53287E+03	0.45562E+00
95	0.10000E+04	0.50000E+03	0.40000E+03	0.53286E+03	0.21103E-01
96	0.10000E+04	0.50000E+03	0.50000E+03	0.53285E+03	0.25960E-01
97	0.10000E+04	0.10000E+04	0.00000E+00	0.53161E+03	0.18427E+01
98	0.10000E+04	0.10000E+04	0.20000E+03	0.53166E+03	0.10197E+01
99	0.10000E+04	0.10000E+04	0.40000E+03	0.53165E+03	0.43592E-01
100	0.10000E+04	0.10000E+04	0.50000E+03	0.53164E+03	0.49508E-01
101	0.10000E+04	0.15000E+04	0.00000E+00	0.52953E+03	0.24778E+01
102	0.10000E+04	0.15000E+04	0.20000E+03	0.52959E+03	0.12279E+01

Table 5.--*Selected output for sample simulation*--Continued

103	0.10000E+04	0.15000E+04	0.40000E+03	0.52958E+03	0.44277E-01
104	0.10000E+04	0.15000E+04	0.50000E+03	0.52957E+03	0.59432E-01
105	0.10000E+04	0.20000E+04	0.00000E+00	0.52661E+03	0.17074E+01
106	0.10000E+04	0.20000E+04	0.20000E+03	0.52665E+03	0.73015E+00
107	0.10000E+04	0.20000E+04	0.40000E+03	0.52664E+03	0.14032E-01
108	0.10000E+04	0.20000E+04	0.50000E+03	0.52664E+03	0.37907E-01
109	0.10000E+04	0.25000E+04	0.00000E+00	0.52303E+03	0.72781E+00
110	0.10000E+04	0.25000E+04	0.20000E+03	0.52306E+03	0.25050E+00
111	0.10000E+04	0.25000E+04	0.40000E+03	0.52304E+03	-0.35839E-02
112	0.10000E+04	0.25000E+04	0.50000E+03	0.52304E+03	0.85501E-02
113	0.10000E+04	0.30000E+04	0.00000E+00	0.51887E+03	0.26513E+00
114	0.10000E+04	0.30000E+04	0.20000E+03	0.51903E+03	0.73789E-01
115	0.10000E+04	0.30000E+04	0.40000E+03	0.51920E+03	0.14296E-02
116	0.10000E+04	0.30000E+04	0.50000E+03	0.51921E+03	0.35718E-02
117	0.10000E+04	0.35000E+04	0.00000E+00	0.51619E+03	0.30831E-01
118	0.10000E+04	0.35000E+04	0.20000E+03	0.51613E+03	0.25067E-02
119	0.10000E+04	0.35000E+04	0.40000E+03	0.51606E+03	-0.15160E-02
120	0.10000E+04	0.35000E+04	0.50000E+03	0.51607E+03	-0.90147E-03
121	0.10000E+04	0.40000E+04	0.00000E+00	0.51425E+03	-0.16723E-02
122	0.10000E+04	0.40000E+04	0.20000E+03	0.51440E+03	-0.14602E-03
123	0.10000E+04	0.40000E+04	0.40000E+03	0.51458E+03	0.19407E-03
124	0.10000E+04	0.40000E+04	0.50000E+03	0.51459E+03	0.10417E-03
125	0.10000E+04	0.45000E+04	0.00000E+00	0.51404E+03	0.49928E-03