

Figure 2.--Dissolved solids in the upper A2 aquifer.

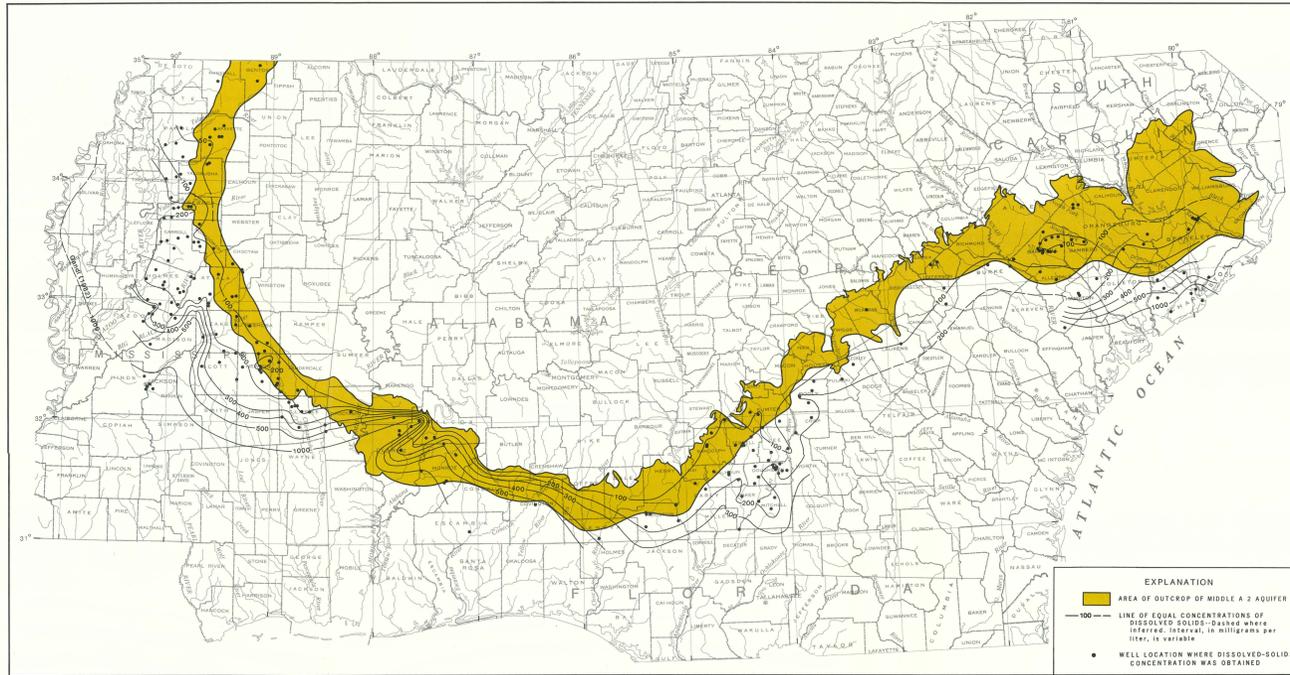


Figure 3.--Dissolved solids in the middle A2 aquifer.

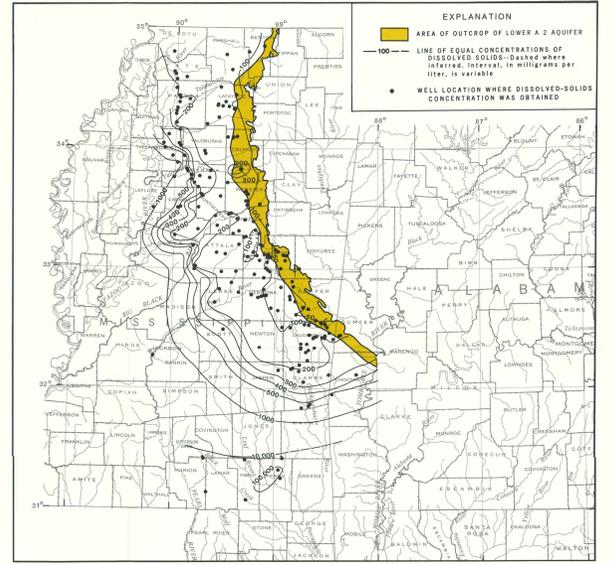


Figure 4.--Dissolved solids in the lower A2 aquifer.

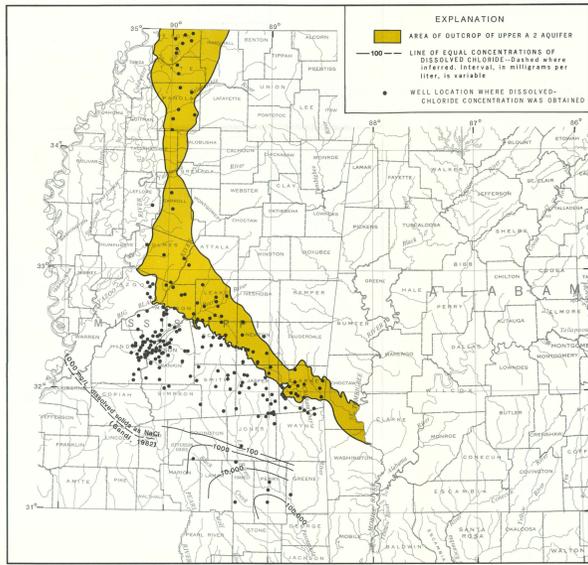


Figure 5.--Dissolved chloride in the upper A2 aquifer.

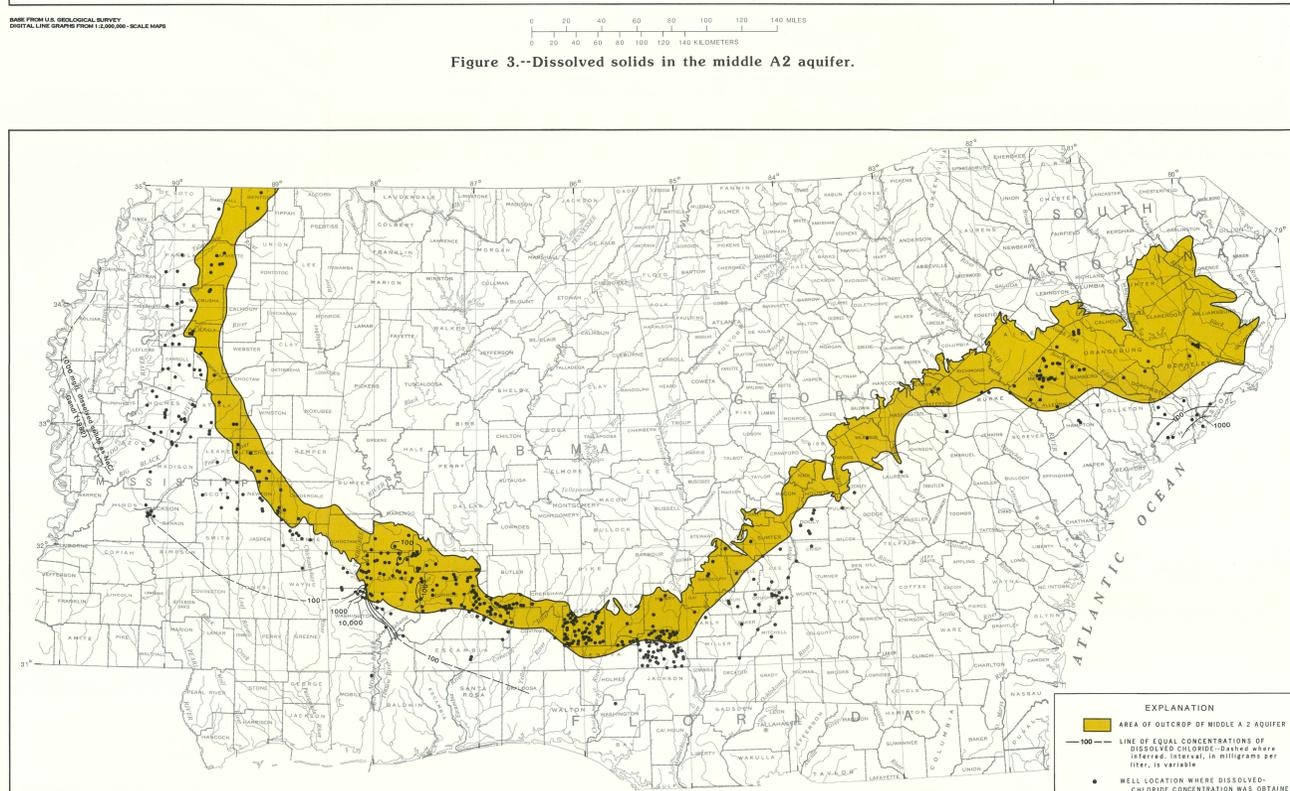


Figure 6.--Dissolved chloride in the middle A2 aquifer.

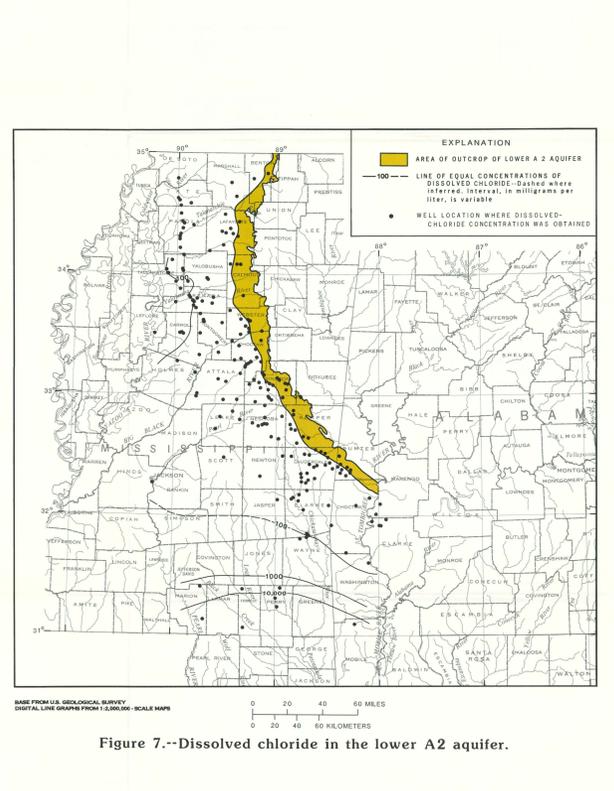


Figure 7.--Dissolved chloride in the lower A2 aquifer.

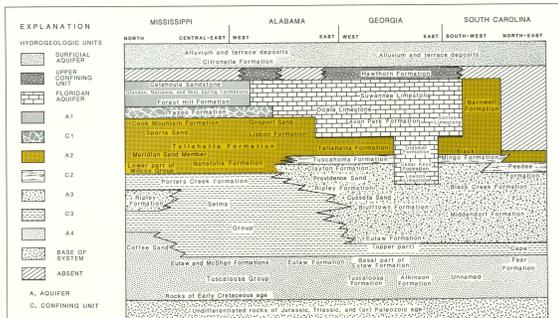


Figure 1.--Schematic diagram of regional hydrogeologic units and selected rock-stratigraphic units in the Southeastern U.S. Coastal Plain. (Modified from Renken, (1984))

**INTRODUCTION**

In the Coastal Plain of Mississippi, Alabama, Georgia, and South Carolina, water-bearing zones comprise major hydrogeologic units of an extensive sand aquifer system that is characterized by regional ground-water flow (Baker, 1985). The aquifer system is being studied as part of the U.S. Geological Survey's Sand Aquifer Regional Aquifer System Analysis (SARASA) program. This system has been delineated by Renken (1984, fig. 1) into regional aquifers and confining beds that correlate with major permeable and impermeable zones in each State within the study area (fig. 2). A series of preliminary reports developed by the SARASA program depict the ground-water geology and water chemistry of the sand aquifers in the Southeastern Coastal Plain.

This report contains maps that depict the spatial distribution of concentrations of dissolved solids, dissolved chloride, dissolved iron, and hydrochemical facies (Bak, 1981) for three water-bearing zones in the A2 aquifer: (1) an upper zone consisting mostly of sands in the Sparta Sand and the Cockfield Formation in Mississippi; (2) a middle zone consisting of sands in the upper part of the Wilcox Group combined with the Meridian Sand Member of the

**WATER-QUALITY MAPS FOR THE MIDDLE TERTIARY AQUIFER IN THE SOUTHEASTERN COASTAL PLAIN OF MISSISSIPPI, ALABAMA, GEORGIA, AND SOUTH CAROLINA**

By  
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**HYDROGEOLOGY**

The hydrogeology of the South Atlantic-Gulf region has been generally described by Cederstrom and others (1979). More detailed descriptions of the geology and hydrology of the Southeastern Coastal Plain may be found in county or multiple county reports in each State. The hydrogeologic framework provides the basis for detailed studies of the regional ground-water flow and supports digital simulations of ground-water flow in the Southeastern Coastal Plain.

The Southeastern Coastal Plain aquifer system has been divided into four regionally extensive aquifers that are separated by confining units consisting of clay, chalk, or other low-permeability materials (Renken, 1984). Each of these regional aquifers can be subdivided into zones and based upon vertical head differences or vertical differences in water chemistry within the aquifer, these differences may be small, or, as occur in the A2 aquifer they may be significant. The differences in head and chemistry may be mutually supportive, or they may not.

The regionally extensive A2 aquifer is composed of sand and shale of Tertiary age which were deposited in deltaic and associated near-shore environments. The aquifer consists of massive to thinly laminated, very fine to coarse, glauconitic, calcareous, quartz sand of late Eocene to middle Paleocene age. In Mississippi, the geology and hydrology of the Tertiary sediments have been previously studied and detailed subsurface information and appropriate water-quality data are available (Bowell, 1976a, 1976b; Newcome, 1976; Spiers, 1976; Wagon, 1985; Gandt, 1982). To better understand the complex water chemistry, the A2 aquifer has been subdivided in Mississippi and western Alabama into three water-bearing zones, the upper, middle, and lower zones. For vertical chemical differences exist in the water chemistry of the A2 aquifer in Alabama, Georgia, and South Carolina. These data closely resemble those from the middle zone in Mississippi and are accordingly mapped as an extension of the zone. The upper and lower zones are mapped separately in Mississippi and western Alabama.

The upper water-bearing zone of the A2 aquifer is limited to Mississippi and western Alabama, and consists of sands in the Sparta Sand, the Cockfield Formation, and the Cook Mountain Formation. The latter formation is less extensive as an aquifer than the previous two, but available chemical data on water from Cook Mountain sediments have been included with the upper water-bearing zone. The Sparta Sand and Cockfield Formation have been previously reported as separate aquifers throughout Mississippi (Newcome, 1976; Spiers, 1976). They are known, however, to be intercorrelated hydrologically in many places, particularly north of Jackson, Mississippi, where the Cook Mountain Formation is sandy and acts as a hydraulic connection between the two aquifers. In deeper parts of the aquifer, dissolved solids locally, but on a regional basis, the chemical data may be combined for these three units.

The quality of ground water from sand aquifers in the Southeastern Coastal Plain is within U.S. Environmental Protection Agency drinking water standards (U.S. Environmental Protection Agency, 1977) in many places. Concentrations of dissolved solids in excess of 500 mg/L may occur as a result of mineral-water interactions in the aquifer, or mixing of fresh ground water with seawater brines present in the deeper parts of the water-bearing zones. Most of the water is sodium and bicarbonate dominated (Lee, 1985) with concentrations of dissolved solids are less than about 500 mg/L. In deeper parts of the aquifer, dissolved sodium and chloride ion dominant (Cederstrom and others, 1979) and concentrations of dis-

solved solids may exceed 10,000 mg/L. Concentrations of dissolved iron may reach 20,000 mg/L in freshwater areas in noncalcareous sands and are usually treated before the water is used.

Maps of concentration of dissolved solids, dissolved iron, dissolved chloride, and hydrochemical facies illustrate spatial distributions of these principal water-quality parameters in the A2 aquifer of the Southeastern Coastal Plain.

Maps of concentration of dissolved solids in the A2 (figs. 2 to 4) aquifer were prepared using values of residue on evaporation at 180°C, where available, or calculated dissolved solids if no residue on evaporation was recorded. Concentrations of dissolved solids in ground water generally increase with distance from the area of outcrop where recharge occurs. In recharge areas, concentration of dissolved solids in ground water may be less than 500 mg/L and near that of dissolved solids observed in rainfall (Jung and Worth, 1968). Waters that have dissolved solids concentrations of less than 500 mg/L usually result from water-rock interactions (Lee, 1985). Concentrations of dissolved solids are greater than 500 mg/L, generally the result of the combined effects of water-rock interactions and mixing of saline water. This mixing produces concentrations of dissolved solids greater than 10,000 mg/L in places. Near Hattiesburg, south-central Mississippi, known salt domes have locally caused elevated concentrations of dissolved solids in ground water (Spiers and Gandt, 1980).

In southern Georgia and parts of South Carolina, ground-water supplies are plentiful and of good quality in the Floridan aquifer system. As a result, water wells generally do not penetrate the deeper sand aquifers and data are not available for the aquifer in this part of the Southeastern Coastal Plain. In Mississippi, contour lines showing the location of 1,000 mg/L dissolved solids as sodium chloride for each water-bearing zone, derived from geophysical data of Gandt (1982), have been used to supplement chemical distribution on the maps.

Ground water in the upper water-bearing zone of the A2 aquifer in Mississippi (fig. 2) increases in concentration of dissolved solids downward, reaching concentrations of greater than 500 mg/L as a result of dissolution of aquifer minerals (Lee, 1985). A closed area of low concentrations of dissolved solids (less than 100 mg/L) is found near the city of Jackson in central Mississippi, and probably reflects vertical differences in concentrations of dissolved solids in the aquifer with most of the low values representing water wells completed in the Cockfield Formation. Dissolved-solids concentrations become greater than 13,000 mg/L south of and Wayne Counties in southwestern Mississippi. Concentrations of dissolved solids exceed 100,000 mg/L in southwestern Alabama.

The distribution of dissolved chloride in the three water-bearing zones shows the extent of freshwater flushing of brines or reflect seawater. In southern Mississippi, elevated chloride concentrations reflect the influence of known salt domes. Mixing of meteoric ground water with saline water is indicated by the pronounced increase in chloride concentration. Further downward, beyond the mixing zone, sparse chemical data show that sodium-chloride brines are present in the deeper parts of the aquifer. Maps of dissolved-chloride concentrations (figs. 5 to 7) show similar trends of increasing concentrations as do dissolved solids where dissolved solids exceed about 500 mg/L.

Concentrations of dissolved chloride in water in the upper zone (fig. 5) are less than 100 mg/L in most areas and are less than 10 mg/L in many places. Concentrations of dissolved chloride exceed 100 mg/L in Hinds and Rankin Counties in central Mississippi and in southern Mississippi and increase to greater than 100,000 mg/L in Perry County in southeastern Mississippi.

Concentrations of dissolved chloride generally are less than 10 mg/L in waters of the middle zone (fig. 6). Near Jackson in central Mississippi, dissolved-chloride concentrations exceed 500 mg/L. Concentrations of dissolved chloride greater than 200 mg/L occur locally in western Wilcox County, Alabama. South of Charleston, South Carolina, dissolved-chloride concentrations exceed 1,000 mg/L, probably as a result of incomplete flushing of seawater from the aquifer. In southeastern Alabama, dissolved-chloride concentrations that exceed 10,000 mg/L may be due to mixing of freshwater with brines present in the aquifer, or may result from upward leakage of saline ground water from deeper aquifers.

Dissolved-chloride concentrations in waters of the lower zone generally are less than 10 mg/L in the upper part of the study area (fig. 7). Concentration increase west and south to greater than 100 mg/L in Citrus and Carroll Counties in northern Mississippi and in Wayne County in southeastern Mississippi. In Forrest and Perry Counties in southern Mississippi, dissolved-chloride concentrations exceed 10,000 mg/L due to salt domes in the area.