

MODIFICATIONS OF A THREE-DIMENSIONAL GROUND-WATER FLOW MODEL TO ACCOUNT  
FOR VARIABLE WATER DENSITY AND EFFECTS OF MULTIAQUIFER WELLS

By A. L. Kontis and R. J. Mandle

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## CONVERSION FACTORS AND ABBREVIATIONS

For readers who prefer metric (International System) units rather than the inch-pound units used in this report, conversion factors are listed below:

<u>Multiply inch-pound unit</u>	<u>By</u>	<u>To obtain metric unit</u>
<u>Length</u>		
inch (in.)	25.40	millimeter (mm)
foot (ft)	0.3048	meter (m)
mile (mi)	1.609	kilometer (km)
<u>Area</u>		
square foot (ft <sup>2</sup> )	0.09290	square meter (m <sup>2</sup> )
square mile (mi <sup>2</sup> )	2.590	square kilometer (km <sup>2</sup> )
<u>Volume</u>		
gallon (gal)	0.003785	cubic meter (m <sup>3</sup> )
million gallons (Mgal)	3,785	cubic meter (m <sup>3</sup> )
cubic foot (ft <sup>3</sup> )	0.02832	cubic meter (m <sup>3</sup> )
<u>Flow</u>		
foot per second (ft/s)	0.3048	meter per second (m/s)
foot per day (ft/d)	0.3048	meter per day (m/d)
cubic foot per second (ft <sup>3</sup> /s)	0.02832	cubic meter per second (m <sup>3</sup> /s)
cubic foot per second per square mile [(ft <sup>3</sup> /s)mi <sup>2</sup> ]	0.01093	cubic meter per second per square kilometer [(m <sup>3</sup> /s)/km <sup>2</sup> ]
million gallons per day (Mgal/d)	0.04381	cubic meter per second (m <sup>3</sup> /s)
<u>Temperature</u>		
degree Fahrenheit (°F)	°C = 5/9 (°F-32)	degrees Celsius (°C)
<u>Hydraulic Conductivity</u>		
foot per second (ft/s)	0.3048	meter per second (m/s)
<u>Transmissivity</u>		
square foot per day (ft <sup>2</sup> /d)	0.09290	square meter per day (m <sup>2</sup> /d)
<u>Pressure</u>		
bar	100	kilopascal (kPa)
pound per square inch (lb/in <sup>2</sup> )	6.895	kilopascal (kPa)
pound per square inch (lb/in <sup>2</sup> )	703.1	kilogram-force per square meter (kg/m <sup>2</sup> )

# CONVERSION FACTORS AND ABBREVIATIONS (continued)

## Dynamic Viscosity

pound-force-second per square foot (lbf.s/ft <sup>2</sup> )	47.88	pascal-second (Pa.s)
slug per foot per second (slug/ft/s)	47.88	pascal-second (Pa.s)

## Acceleration

foot per second squared (ft/s <sup>2</sup> )	0.3048	meter per second squared (m/s <sup>2</sup> )
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## Density

slug per cubic foot (slug/ft <sup>3</sup> )	515.4	kilogram per cubic meter (kg/m <sup>3</sup> )
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## EQUIVALENTS

	<u>Inch-pound unit</u>	<u>Metric unit</u>
Acceleration of gravity	32.17 ft/s <sup>2</sup>	9.806 m/s <sup>2</sup>
Density of pure water (68°F, 20°C)	1.937 slug/ft <sup>3</sup>	998.2 kg/m <sup>3</sup>
Specific weight of water (58°F, 15°C)	62.4 lb/ft <sup>3</sup>	9,800 n/m <sup>3</sup>
Standard atmosphere	14.70 lb/in <sup>2</sup>	101,325 kPa

Sea Level: In this report "sea level" refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929)--a geodetic datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called "Sea level datum of 1929."

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## ABSTRACT

A three-dimensional finite-difference ground-water-flow model was modified and used in the U.S. Geological Survey's Northern Midwest Regional Aquifer System Analysis to simulate the effects of wells open to various combinations of aquifers and the effects of ground-water density that varies spatially but is time invariant. The computer code for the slice-successive-over-relaxation solution method used to solve for head and a description of how it functions is given, as well as modifications that were made to account for the hydraulic effects of multi-aquifer-wells. In addition, two programs are presented that compute data required for the simulation of variable-density effects: one program calculates model-block conductances and variable-density terms external to the ground-water-flow model; the other computes ground-water and pure-water density and viscosity for input to the first program.

## INTRODUCTION

As part of the U.S. Geological Survey's Regional Aquifer-System Analysis (RASA) program (Bennett, 1979; Steinhilber and Young, 1979; Young and others, 1986), a regional ground-water flow simulation (Mandle and Kontis, in press) of the Cambrian-Ordovician aquifer system of the northern Midwest was developed through a modified version of the three-dimensional finite difference ground-water-flow model (Trescott, 1975; Trescott and Larson, 1976). This model is referred to herein as the Trescott-Larson model. A 378,880-mi<sup>2</sup> area was modeled that includes large areas east and south of the study area (fig. 1) to satisfy boundary conditions.

The effects of two conditions in the model area--the presence of brines in the Illinois and Michigan basins, and the presence of wells open to more than one aquifer--could not be simulated with the documented version of the Trescott-Larson model because it assumes that ground water is of constant density and that a given well is open only to a single aquifer.

To simulate the effect of brines of variable density on the regional flow system, an approach suggested by Bennett (1980) was used, which assumes that, although temperature and solute concentration may vary spatially, they are invariant over the simulated period (120 years). Under this assumption, the finite-difference ground-water-flow equation can be written to accommodate the effects of variations in water density such that the Trescott-Larson model program can be used with only minor modifications to the computer code. This is done by precomputing model hydraulic conductances and water-density terms through a computer program termed VARDEN (VARIABLE DENSITY), which is separate from the Trescott-Larson model.

The majority of wells in the study area are open to more than one aquifer except in Minnesota (fig. 1). As discussed by Bennett and others (1982), multiaquifer wells can markedly affect the nature of flow between aquifers connected by the wells. The flow equation that is solved in the Trescott-Larson model links the head at a given node with only the heads in the four adjacent nodes within a model layer and with the two adjacent nodes in the layers above and below the node. For a well open to more than one aquifer, internal flow between the aquifers through the open well is not accounted for. In addition, most measured water levels in multiaquifer wells that are available for model calibration are a composite of generally unequal heads. Consequently, to simulate the effects of multiaquifer wells and to provide a means of calibrating model results, additional modifications were made to the flow equation and to the Trescott-Larson model code. To implement the modified finite-difference flow equation, the slice-successive-over-relaxation (SSOR) solution scheme was used.

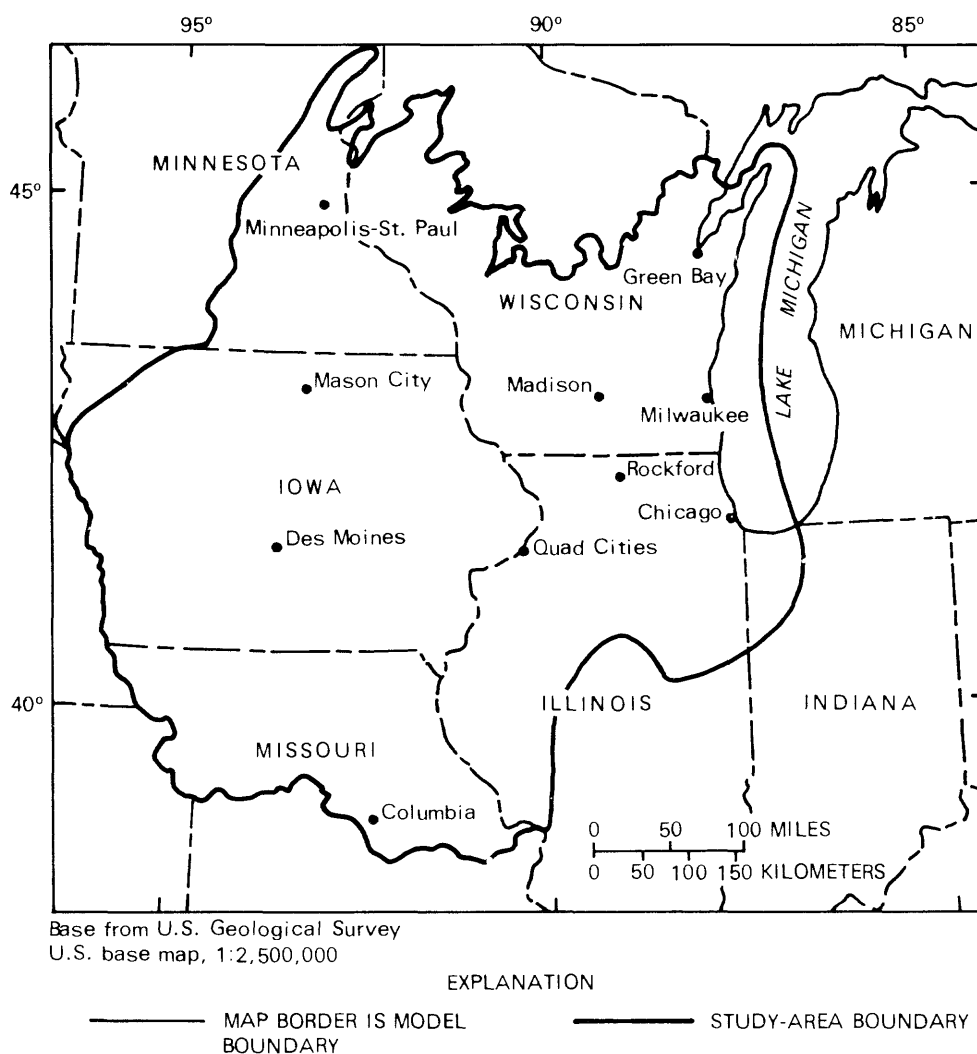


Figure 1.--Northern Midwest Regional Aquifer Systems Analysis study area and boundary of ground-water flow model.



## Purpose and Scope

The development of the modified flow equation and results of the northern Midwest ground-water flow simulation are discussed in detail in Mandle and Kontis (in press). This report (1) presents the computer modifications of the Trescott-Larson model and SSOR algorithm with which variable-density and multiaquifer-well effects were simulated, and (2) describes (a) the VARDEN program, which precomputes model hydraulic conductances and variable-density terms, and (b) the STEAM program, which computes ground-water density, pure-water density, and viscosity for input to VARDEN. The original SSOR-solution computer code for use with the Trescott-Larson model was written by S. P. Larson and C. R. Faust (U.S. Geological Survey, written commun., 1979). Because the original code was not formally documented, it is given here in its entirety (appendix F) with a general explanation of how the subroutine functions (appendix A). In addition, the variable-density method is compared with results from a cross-sectional electric analog model of an area in Puerto Rico (Bennett and Guisti, 1971).

The purpose of this report is to present the modifications and support programs in sufficient detail that the computational aspects of the variable-density and multiaquifer-well effects used in the northern Midwest regional simulation (Mandle and Kontis, in press) are clearly understood. The modifications and programs were developed specifically for the northern Midwest study, but the variable-density and multiaquifer-well formulations could be used for other applications as well. Although the report is not intended to be a users manual, the text and comment statements in program listings give sufficient detail that the procedures can be readily adapted to systems with similar hydrogeologic conditions. The reader is assumed to be familiar with the Trescott-Larson model or to have access to the original documentation (Trescott, 1975; Trescott and Larson, 1976). Where possible, this report uses the same notation as was used in the original documentation.

## MODIFICATIONS

### Variable Water Density

#### Finite-Difference Equation

As shown in Mandle and Kontis (in press), the quasi-three-dimensional finite-difference equation describing ground water flow in terms of freshwater head,  $h_f$ , for spatially varying ground-water density, may be expressed as:

$$\begin{aligned} & -TR_{i,j-1,k}(h_{f,j,k} - h_{f,j-1,k}) + A_{i,j-1,k} + TR_{i,j,k}(h_{f,j+1,k} - h_{f,j,k}) - A_{i,j,k} \\ & - TC_{i-1,j,k}(h_{f,j,k} - h_{f,i-1,j,k}) + B_{i-1,j,k} + TC_{i,j,k}(h_{f,i+1,j,k} - h_{f,i,j,k}) - B_{i,j,k} \\ & - TK_{i,j,k-1}(h_{f,j,k} - h_{f,j,k-1}) + \Gamma_{i,j,k-1} + TK_{i,j,k}(h_{f,j,k+1} - h_{f,j,k}) - \Gamma_{i,j,k} \\ & = S^* \frac{\Delta h_{f,i,j,k}}{\Delta t} - Q^* \end{aligned} \quad (1)$$

where:

$TR_{i,j-1,k}$  and  $TR_{i,j,k}$  are the aquifer-layer conductances along rows between pairs of nodes  $(i,j-1,k)$ ,  $(i,j,k)$  and  $(i,j,k)$ ,  $(i,j+1,k)$  (fig. 2C);

$TC_{i-1,j,k}$  and  $TC_{i,j,k}$  are the aquifer-layer conductances along columns between pairs of nodes  $(i-1,j,k)$ ,  $(i,j,k)$  and  $(i,j,k)$ ,  $(i+1,j,k)$  (fig. 2B); and

$TK_{i,j,k}$  and  $TK_{i,j,k-1}$  are conductances of the confining layer between pairs of aquifer layer nodes  $(i,j,k+1)$ ,  $(i,j,k)$  and  $(i,j,k)$ ,  $(i,j,k-1)$  (fig. 2B and 2C).

The nodes are taken to be at the center of model blocks defined to be the mid-point between the top and bottom surfaces of aquifer layers. In equation 1, unscripted terms are understood to be at node  $i,j,k$ .

Equation 1 is developed in Mandle and Kontis (in press) by applying the variable-density form of Darcy's law and balancing the mass of equivalent pure water entering and leaving an aquifer element, on the assumption that the dispersion of pure water is negligible. An alternative development of variable density effects, based on the conservation of pure-water mass, is given in Weiss (1982a). A method based on conservation of ground-water mass is described by Kuiper (1983, 1985).

In figure 2, the nodes are depicted as dipping along an  $\alpha$  axis inclined  $a'$  degrees with respect to the  $x$  (horizontal) axis of an  $x,y,z$  cartesian coordinate system (fig. 2C) and along a  $\beta$  axis inclined  $b'$  degrees with respect to the  $y$  (horizontal) axis (fig. 2B). The normal to the plane of  $\alpha$  and  $\beta$  defines a  $\gamma$  axis, which is at an angle  $c'$  with respect to the  $z$  (vertical) axis (fig. 2A).

In equation 1, the  $A$ ,  $B$ , and  $\Gamma$  terms represent the effects of gravity forces in the  $\alpha$ ,  $\beta$ , and  $\gamma$  directions arising from the presence of variable-density ground water. They are computed from:

$$A_{i,j-1,k} = -TR_{i,j-1,k}(\rho' - 1)_{i,j-1/2,k}(Z_{i,j,k} - Z_{i,j-1,k}) \quad (2a)$$

$$A_{i,j,k} = -TR_{i,j,k}(\rho' - 1)_{i,j+1/2,k}(Z_{i,j+1,k} - Z_{i,j,k}) \quad (2b)$$

$$B_{i-1,j,k} = -TC_{i-1,j,k}(\rho' - 1)_{i-1/2,j,k}(Z_{i,j,k} - Z_{i-1,j,k}) \quad (2c)$$

$$B_{i,j,k} = -TC_{i,j,k}(\rho' - 1)_{i+1/2,j,k}(Z_{i+1,j,k} - Z_{i,j,k}) \quad (2d)$$

$$\Gamma_{i,j,k-1} = -TK_{i,j,k-1}b_L(\rho' - 1)_{i,j,k-1/2}\left(\frac{\Delta x \Delta y}{\Delta \alpha \Delta \beta}\right)_{i,j,k} \quad (2e)$$

$$\Gamma_{i,j,k} = -TK_{i,j,k}b_L(\rho' - 1)_{i,j,k+1/2}\left(\frac{\Delta x \Delta y}{\Delta \alpha \Delta \beta}\right)_{i,j,k} \quad (2f)$$

where  $\rho'$  is the relative density defined to be the ratio of the density of ground water of interest,  $\rho_s$ , to the density of the equivalent freshwater,  $\rho_f$ , at a temperature of 20 °C. The ground-water density ( $\rho_s$ ) is equal to the combined densities of solvent (pure water) and solutes in the solvent. Relative density at the finite-difference block faces is computed by averaging the densities of adjacent nodes. For example at the  $i, j + 1/2, k$  block face,

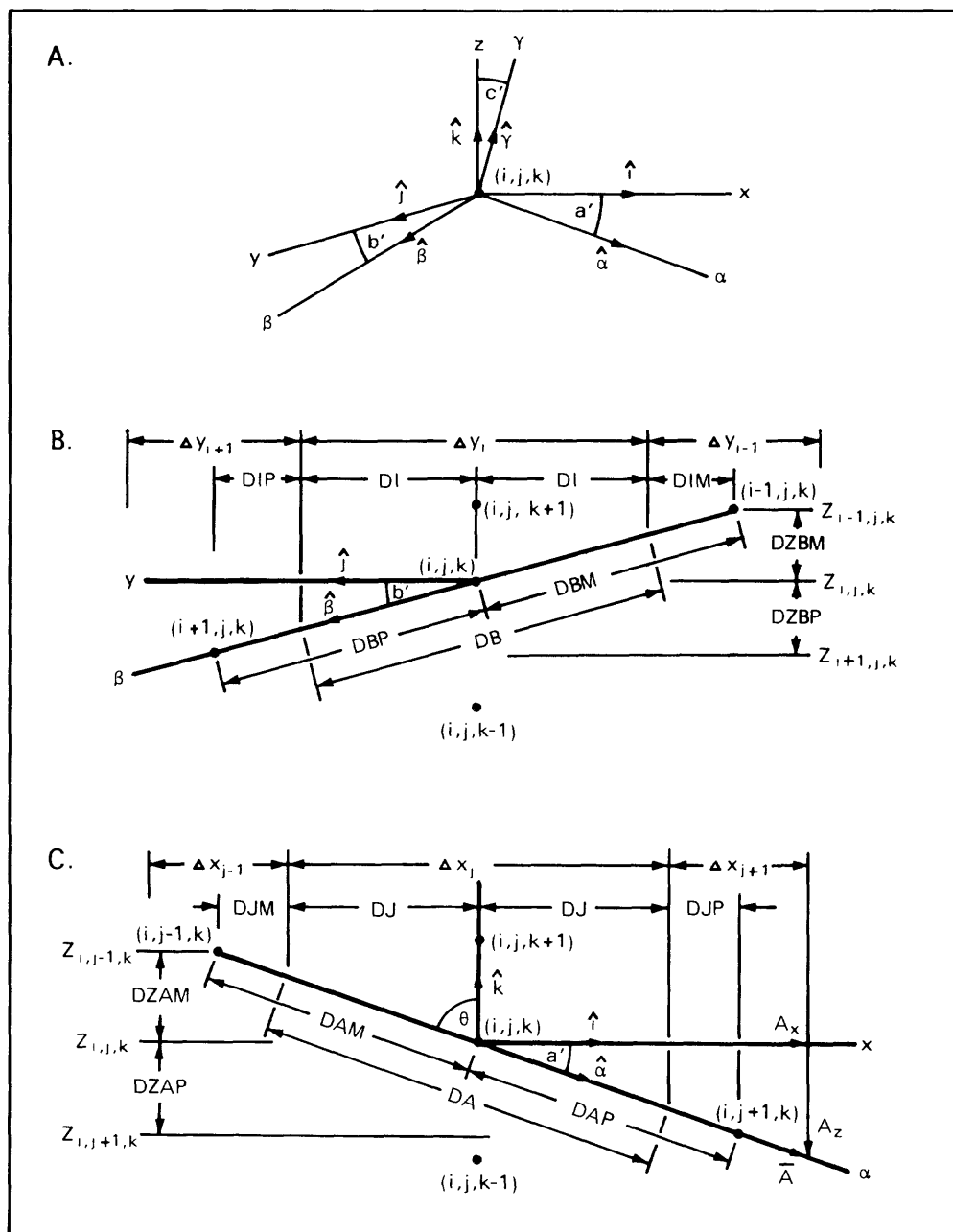


Figure 2.--(A) Variable density  $\alpha, \beta, \gamma$  coordinate system relative to  $x, y, z$  coordinate system, and geometric terms used in program VARDEN in (B),  $y, \beta$  and (C),  $x, \alpha$  planes.

$$\rho'_{i,j+1/2,k} = \frac{\rho_{s_{i,j,k}} + \rho_{s_{i,j+1,k}}}{2\rho_f}.$$

The Z terms of equations 2a-d are the differences in elevation between adjacent nodes in the  $\alpha$  and  $\beta$  directions. The term  $\frac{\Delta x}{\Delta \alpha} \frac{\Delta y}{\Delta \beta}$  of equations 2e and

2f, where  $\Delta x$  and  $\Delta y$  are the horizontal dimensions of the finite-difference block i,j,k of the aquifer, approximates the gradient of elevation in the direction normal to the plane of the aquifer (fig.2A). A derivation of this approximation is given in the discussion of program VARDEN (p. 13).

The conductance terms of equations 1 and 2 formed from the distance-weighted harmonic mean of hydraulic properties of adjacent nodes, are given by:

$$\begin{aligned} TR_{i,j-1,k} &= \frac{\Delta \beta \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j-1/2,k}}{\alpha_{i,j,k} - \alpha_{i,j-1,k}} \\ &= \frac{2\Delta \beta \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j,k} \cdot \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j-1,k}}{\Delta \alpha_{i,j,k} \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j-1,k} + \Delta \alpha_{i,j-1,k} \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j,k}}. \end{aligned} \quad (3a)$$

$$TR_{i,j,k} = \frac{2\Delta \beta \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j,k} \cdot \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j+1,k}}{\Delta \alpha_{i,j,k} \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j+1,k} + \Delta \alpha_{i,j+1,k} \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j,k}}, \quad (3b)$$

$$TC_{i-1,j,k} = \frac{2\Delta \alpha \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i,j,k} \cdot \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i-1,j,k}}{\Delta \beta_{i,j,k} \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i-1,j,k} + \Delta \beta_{i-1,j,k} \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i,j,k}}, \quad (3c)$$

$$TC_{i,j,k} = \frac{2\Delta \alpha \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i,j,k} \cdot \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i+1,j,k}}{\Delta \beta_{i,j,k} \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i+1,j,k} + \Delta \beta_{i+1,j,k} \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i,j,k}}, \quad (3d)$$

$$TK_{i,j,k-1} = \Delta \alpha \Delta \beta \left( \frac{K_{\gamma\gamma} \rho_p}{b_t \mu'} \right)_{i,j,k-1/2}, \text{ and} \quad (3e)$$

$$TK_{i,j,k} = \Delta \alpha \Delta \beta \left( \frac{K_{\gamma\gamma} \rho_p}{b_t \mu'} \right)_{i,j,k+1/2} \quad (3f)$$

In equation set 3,  $K_{\alpha\alpha}$ ,  $K_{\beta\beta}$ ,  $K_{\gamma\gamma}$  are the components of the hydraulic-conductivity tensor in the  $\alpha$ ,  $\beta$ ,  $\gamma$  direction (L/T) of the equivalent freshwater aquifer having ground water of density  $\rho_f$  and dynamic viscosity of  $\mu_f$ , and

$\rho_p$  = pure-water density of node  $i,j,k$  defined to be the difference between ground-water density,  $\rho_s$ , and the mass of solute per unit volume of solution;

$\mu'$  = relative viscosity, defined to be the ratio of the dynamic viscosity of the ground water of interest ( $\mu$ ) to the dynamic viscosity ( $\mu_f$ ) of the equivalent freshwater at 20°C (dimensionless);

$b$  = average thickness of aquifer layer node  $i,j,k$  (L)

$\Delta\alpha, \Delta\beta$  = inclined dimensions of the finite-difference block of the aquifer, (L) (fig.2); and

$b_U, b_L$  = thickness of confining layers above and below node  $i,j,k$  (L).

As discussed in Mandle and Kontis (in press), the variable-density formulation is based on the assumption that ground-water density varies spatially but is time invariant over the period of simulation. If pumping centers are far from areas of variable density, changes in head over time are assumed to not significantly alter the density and viscosity distribution; therefore the formulation (eq. 1) may be applied for transient simulation. Accordingly, on the right-hand side of equation 1:

$S^*$  = modified storage coefficient and is the product of the pure-water density ( $\rho_p$ ), storage coefficient ( $S$ ), and block area ( $\Delta\alpha\Delta\beta$ );

$\Delta h_f$  = change in freshwater head per time increment ( $\Delta t$ );

$Q^*$  = product of pure-water density and  $Q$ ; a source (positive) or sink (negative) term representing addition or removal of water from the node.

Thus:

$$S^* = (\Delta\alpha\Delta\beta\rho_p S)_{i,j,k}, \text{ and}$$

$$Q^* = (\rho_p Q)_{i,j,k}. \quad (4)$$

In equation 1, the variable-density terms given by equations 2a-f are known and are independent of head so that they can be incorporated into a single term  $(TS_{i,j,k})$ . Thus, equation 1 can be rewritten as:

$$\begin{aligned} & -TR_{i,j-1,k}(h_{i,j,k} - h_{i,j-1,k}) + TR_{i,j,k}(h_{i,j+1,k} - h_{i,j,k}) \\ & -TC_{i-1,j,k}(h_{i,j,k} - h_{i-1,j,k}) + TC_{i,j,k}(h_{i+1,j,k} - h_{i,j,k}) \\ & -TK_{i,j,k-1}(h_{i,j,k} - h_{i,j,k-1}) + TK_{i,j,k}(h_{i,j,k+1} - h_{i,j,k}) \\ & = S^* \frac{\Delta h_{f,i,k}}{\Delta t} - Q^* - TS_{i,j,k}, \end{aligned} \quad (5a)$$

where

$$TS_{i,j,k} = A_{i,j-1,k} - A_{i,j,k} + B_{i-1,j,k} - B_{i,j,k} + \Gamma_{i,j,k-1} - \Gamma_{i,j,k} . \quad (5b)$$

Also

$$\begin{aligned} PSA_{i,j,k} &= + A_{i,j,k} - A_{i,j-1,k} , \\ PSB_{i,j,k} &= + B_{i,j,k} - B_{i-1,j,k} , \text{ and} \\ PS\Gamma_{i,j,k} &= + \Gamma_{i,j,k} - \Gamma_{i,j,k-1} . \end{aligned} \quad (5c)$$

The form of equation 5a is the same as that of the finite-difference equation of the Trescott-Larson model except for the TS term. This term represents the net effect of gravity forces exerted on the flow by the mass of ground water in a tilted aquifer block.

In areas containing relatively fresh water, the individual variable-density terms A, B, and  $\Gamma$  (eq. 2a-f) are zero or very small so that the magnitude of flow across model block faces is proportional to freshwater-head gradient, and flow is essentially orthogonal to contours of freshwater head. In areas containing brines, and where aquifers are not horizontal, the flow may be significantly altered by the A, B, and  $\Gamma$  terms, and resultant flow directions will not be orthogonal to lines of equal freshwater head. The condition for no flow in these areas is that the freshwater-head gradient in each direction be exactly opposed by the effect of gravity forces in that direction.

The volumetric flow components for any node i,j,k are computed by:

$$QR_{i,j,k} = [-TR_{i,j,k}(h_{f,i,j+1,k} - h_{f,i,j,k}) + A_{i,j,k}]/\rho_{p,i,j+1/2,k} , \quad (6a)$$

$$QC_{i,j,k} = [-TC_{i,j,k}(h_{f,i+1,j,k} - h_{f,i,j,k}) + B_{i,j,k}]/\rho_{p,i+1/2,j,k} , \text{ and} \quad (6b)$$

$$QK_{i,j,k} = [-TK_{i,j,k}(h_{f,i,j,k+1} - h_{f,i,j,k}) + \Gamma_{i,j,k}]/\rho_{p,i,j,k+1/2} , \quad (6c)$$

where:

$QR_{i,j,k}$  = flow along rows between nodes i,j,k and i,j+1,k ( $L^3/T$ )

$QC_{i,j,k}$  = flow along columns between nodes i,j,k and i+1,j,k ( $L^3/T$ ), and

$QK_{i,j,k}$  = flow across confining layers between nodes i,j,k and i,j,k + 1 ( $L^3/T$ ), and

the pure-water density term in each expression is the average value for the pair of nodes in question.

The variable-density finite-difference equation in terms of pressure (P) can be readily obtained from the relation  $h_f = \frac{P}{\rho_{fg}} + Z$ . The conversion to a

pressure formulation may be implemented by (1) replacing all terms in equations 1, 5, and 6 containing  $h_f$  by P, (2) dividing all conductances (eq. 3a-f) by  $\rho_{fg}$ , (3) replacing all terms containing  $(\rho'-1)$  in equations 2a-f by  $\rho'$ , and (4) multiplying all terms containing  $\rho'$  by  $\rho_{fg}$ .

The finite-difference variable-density flow formulation (eq. 5a) was tested by comparing model results with the results of a cross-sectional electric analog model of a freshwater/saltwater-interface problem discussed by Bennett and Giusti (1971). Results of the comparison are shown in figure 3, which depicts the simulated north-south steady-state flow system of an alluvial water-table aquifer along the south-central coast of Puerto Rico. In the analog simulation of Bennett and Giusti, the northern boundary (figs. 3B, 3C) and upper boundary (fig. 3A) were treated as constant heads. The northern constant-head boundary represents uniform flow from the north, and the upper constant-head boundary represents a rainy-season, high-water-table condition north of the coast and heads within the seabed seaward of the coast. The bottom of the model (bedrock) and a freshwater/saltwater interface to the south were treated as zero-flow boundaries. With this configuration and a horizontal hydraulic conductivity of 33 ft/d for alluvium and a vertical anisotropy in conductivity of 1,000, Bennett and Giusti obtained a head pattern shown by the dotted contours in figure 3B.

To test the finite-difference variable-density formulation, the cross section was discretized into 11 layers, each with a thickness of 50 feet (ft) and a uniform spacing of 800 ft along each layer. The layers were assumed to dip uniformly seaward at an angle of  $0.25^\circ$  (23 ft per mile). Except for the freshwater/saltwater interface, all boundaries were treated the same as in the analog model. A profile of the constant-head upper layer is shown in figure 3A. For the variable-density formulation, pure water density,  $\rho_p$ , and relative viscosity,  $\mu'$ , were taken as unity in equations 3a-f and the density of salty ground water (eq. 2a-f) was set at 1.025 grams per cubic centimeter ( $\text{g/cm}^3$ ).

The freshwater heads obtained with the finite-difference model are shown in figure 3B, and a representation of the flow is given in figure 3C. In figure 3C, the tail of each vector is at a model node, and the length and direction of each vector is the resultant of flow within the layer (eq. 6a) and flow across the layer (eq. 6c). The largest vector has a magnitude of 0.026 cubic feet per second ( $\text{ft}^3/\text{s}$ ), and the length of all other vectors is proportional to this maximum length. The magnitude is also depicted by the dotted lines of equal flow in figure 3C. Seaward of the freshwater/saltwater interface, the magnitude of flow is quite small (less than  $0.005 \text{ ft}^3/\text{s}$ ), and the direction of flow is essentially random. Within this region, the freshwater-head difference between nodes of adjacent layers (fig. 3B) is about 1.3 ft. The hydraulic conductance between layers is 0.00941 (slug/s)/ft (slug per second per foot,) and the  $\Gamma$  source term (eq. 2f) is -0.012 slug/s. Consequently, from equation 6c, the flow between layers (QK) in this region is  $(1.3 \times 0.0094 - 0.012)$  divided by the density of pure water--essentially zero flow.

#### Characteristics of Water-Density Terms

The effects of local variations in water density on the simulated ground-water flow system are manifested directly through the term  $TS_{i,j,k}$  of equation 5a. Depending on the sign of this term, a quantity of water is, in effect, added or removed from each model node containing a relative ground-water density,  $\rho'$ , other than unity. Because of this, the TS term may be referred to as a total-source term and A,B, $\Gamma$  (eq. 2a-f) and PSA, PSB and PS3 (eq. 5c) as individual or component-source terms. The magnitude and sign of the total-source term depends on the relative magnitudes of hydraulic conduc-

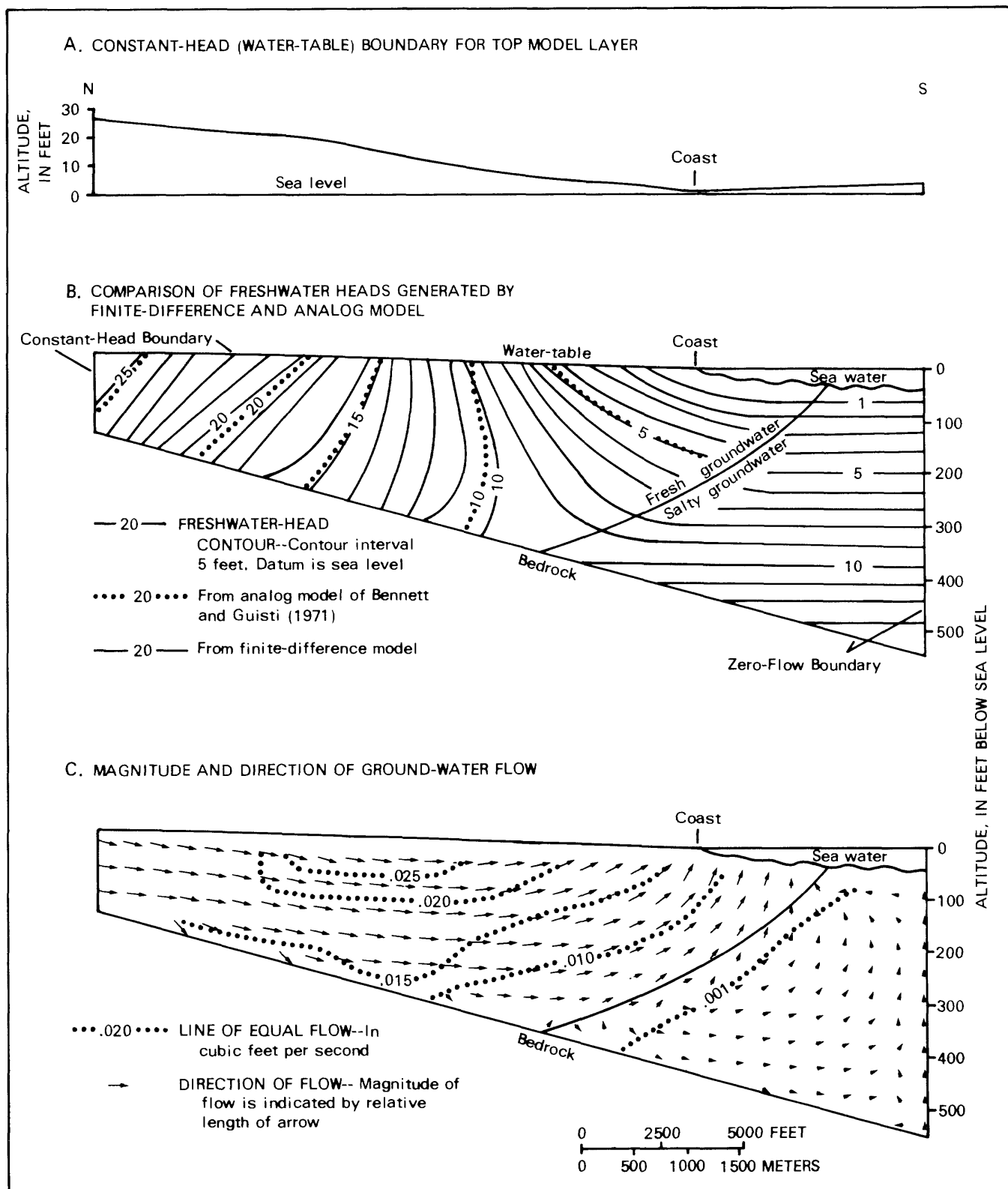


Figure 3.--Results of variable-density finite-difference cross-sectional model of an area near Ponce, Puerto Rico.



tivity, pure-water and ground-water density, viscosity, aquifer and confining-layer thickness, and elevation gradient between the node center and the adjacent node centers according to equations 3a-f and equations 2a-f. Consequently, whether or not the total-source term for a particular node represents a source or sink is difficult to determine from cursory inspection of the data.

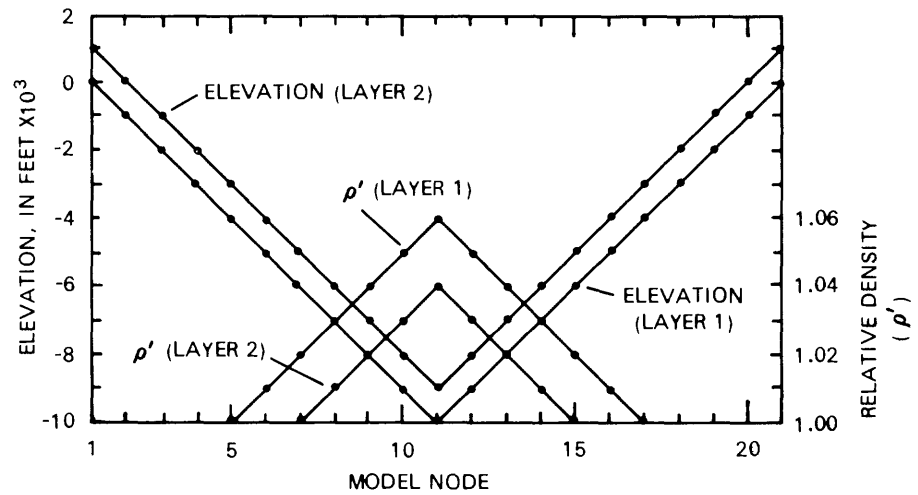
The disposition of the component source terms A, B, and  $\Gamma$  of equations 2a-f used to compute flow in the  $\alpha$ ,  $\beta$ , and  $\gamma$  directions is more readily apparent, however. In areas where the relative ground-water density is greater than unity,  $(\rho'-1)$  is positive, so that the sign of the individual source terms A and B governing flow within a layer is given by the elevation gradient in the appropriate direction. For example, if the elevation of node  $i,j+1,k$  is less than that of node  $i,j,k$ , then  $A_{i,j,k}$  is always positive. In addition, as shown in equation 2f, areas where relative ground-water density is greater than unity will always have a negative  $\Gamma$  term.

To illustrate the nature of the total and component-source terms, their values for a simple example consisting of a cross section through an idealized, v-shaped, two-dimensional basin are shown in figures 4 and 5. The basin consists of three aquifer layers separated by confining layers. The layers have a constant elevation gradient of 1000 ft per model node spacing (fig. 4A). Layers 1 and 2 contain ground water whose relative density is unity at and near the basin margins but increases linearly toward the basin trough. The relative density of the topmost layer (layer 3) is that of freshwater. The hydraulic conductance (TR) in the  $\alpha$  direction (perpendicular to the trough) is set at 0.001 (slug/ft)/s for all nodes of layers 1 and 2. With these values, the A source terms for layer 1 and 2 have the form shown in figure 4B. The A source term for both layers is positive along the descending limb (negative elevation gradient) and negative along the ascending limb (positive elevation gradient).

For the  $\Gamma$  term, the product of all terms in equations 2e and 2f, except for the  $(\rho'-1)$  term, is assumed to be 1 slug/s between the nodes of layers 1 and 2 and layers 2 and 3. For these values and the density distribution of figure 4A, the  $\Gamma$  source terms for layers 1 and 2 have the form shown in figure 4B. Because layer 3 contains freshwater,  $(\rho'-1)$  is zero, making the A term zero for all nodes. In addition, layer 3 has no  $\Gamma$  term because it is the topmost layer.

The components of the total-source terms (PSA and PS3) in the  $\alpha$  and  $\gamma$  directions according to equation 5c are given in figure 5A and 5B. In a simulation of the system, the actual input to the model for each node would consist of the total-source term, TS, obtained by summing these components over each node by equation 5B. As shown in figure 5A, the total-source terms for layer 1 consist primarily of positive contributions with a relatively large positive source at the trough of the basin. The total-source terms for layer 2 (fig. 5B) are negative except for the node at the trough. The total sources for layer 3, equivalent to the  $\Gamma$  terms of layer 2 (fig. 4B) (and not shown on fig. 5), are negative. In a simulation, the relatively large positive total sources of layers 1 and 2 at node 11 (fig. 5) would cause a buildup of freshwater head at the center of the basin. Note that the sum of the total-source terms over all layers (TS of fig. 5A (+0.28 slug/s) + TS of fig. 5B (-0.20 slug/s) +  $\Gamma$  of layer 2 in fig. 4B (-0.08 slug/s)) is zero, as required by conservation of mass.

A. ELEVATION PROFILE, LAYER 2



B. MAGNITUDE OF VARIABLE-DENSITY SOURCE

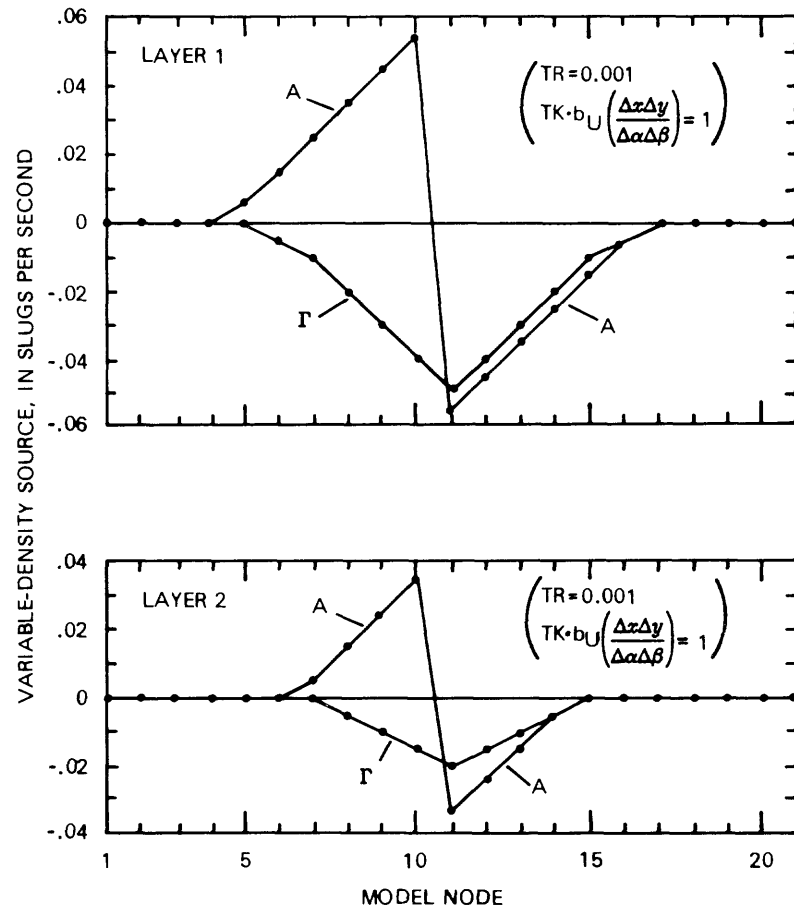
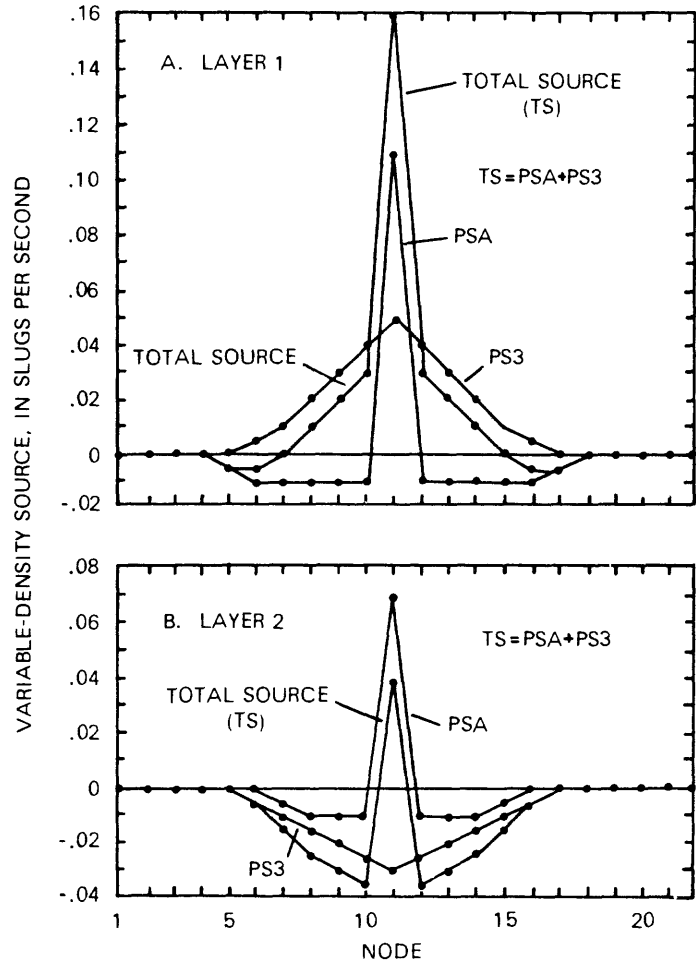


Figure 4.--Profiles of (A) elevation and relative ground-water density, and (B) magnitude of component source terms (A and  $\Gamma$ ) for idealized two-dimensional basin.

Figure 5.

Magnitude of variable water-density source terms (PSA, PS3, and TS) for idealized two-dimensional basin.



#### Description of Program VARDEN

To obtain the variable-density source terms and conductances required for input to the Trescott-Larson finite-difference model, in terms of either freshwater head or pressure, a Fortran computer program, VARDEN (VARIABLE DENSITY), was developed from a program originally written by Weiss (1982a). VARDEN computes and stores exterior to the Trescott-Larson model, the TR, TC, TK conductances (eq. 3a-f) and the modified storage coefficient  $S^*$  (eq. 4). In addition, the individual density terms A, B, and  $\Gamma$  (eqs. 2a-f) for use in postmodel flow calculations (eq. 6a-c) and a modified transmissivity for use in the simulation of the effects of multiaquifer wells, discussed in a later section are computed and stored. For isotropic conditions, the modified transmissivity is:

$$T^*_{i,j,k} = \left( \frac{K_{\alpha\alpha} b \rho_p}{\mu'} \right)_{i,j,k} = \left( \frac{K_{\beta\beta} b \rho_p}{\mu'} \right)_{i,j,k}, \quad (7)$$

The input to the program consists of a single control record containing the size of the model and options to be used, followed by a sequence of arrays (table 1A). For these input arrays, table 1 gives the program names, the corresponding names of equations 1-7, and the number of arrays required for a model with  $K_0$  layers. All arrays, except model-grid spacing, are presumed to reside in a single file (unit LIN). The input transmissivity and confining

Table 1.--Program VARDEN (VARIABLE DENSITY) Input and Output Arrays

CONTROL RECORD IN 615 FORMAT						
Name in program		Description				
IO	Number of model rows					
JO	Number of model columns					
KO	Number of model layers					
LIN	Fortran unit number for input arrays					
IPRESS =	{ 0 for freshwater formulation 1 for pressure formulation					
IDELTA =	{ 1 if confining layer conductivity is to be modified 0 otherwise					
A. INPUT ARRAYS						
Name in program	Description	Name in Eq. 1-7	Number of arrays	Fortran unit number	Format	Units
S	Storage coefficient	S	KO	LIN	10E8.2	--
T	Transmissivity	$K_{\alpha\alpha}^*b, K_{\beta\beta}^*b$	KO	LIN	10E8.2	ft <sup>2</sup> /s
ELEV	Elevation of nodes	Z	KO	LIN	10F8.0	ft
DEN	Ground water density	$\rho_s$	KO	LIN	10F8.1	kg/m <sup>3</sup>
DENP	Pure water density	$\rho_p$	KO	LIN	10F8.1	kg/m <sup>3</sup>
TKV	Confining layer leakance	TK	KO-1	LIN	10E8.2	1/s
THK	Confining layer thickness	$b_U, b_L$	KO-1	LIN	10F8.0	ft
VISC	Viscosity	$\mu$	KO	LIN	10F8.3	cp
DELTA	Factors to modify T and TK	--	KO	LIN	10F8.3	--
DELX	Node spacing along rows	$\Delta x$	1	5	8G10.0	ft
DELY	Node spacing along columns	$\Delta y$	1	5	8G10.0	ft
B. OUTPUT ARRAYS						
Name in program	Description	Name in Eq. 1-7	Number of arrays	Fortran unit number	Format	Units
Arrays for input to Trescott model						
SOUT	Modified storage coefficient	$S^*$	KO	19	10F8.0	slug/ft
TOUT	Modified transmissivity	$K^*b\rho_p/\mu'$	KO	18	8E10.4	(slug/s)/ft
TK	Conductance between layers	TK	KO-1	17	8E10.4	(slug/s)/ft
PST	Total source term	TS	KO	16	8E10.4	slug/s
TR	Conductance along rows	TR	KO	11	8E10.4	(slug/s)/ft
TC	Conductance along columns	TC	KO	12	8E10.4	(slug/s)/ft
Arrays for postmodel computation of flow						
ALPHA	Source term for rows	A	KO	13	8E10.4	slug/s
BETA	Source term for columns	B	KO	14	8E10.4	slug/s
GAMMA	Source term for layers	$\Gamma$	KO-1	15	8E10.4	slug/s

layer leakance arrays are estimates of the hydraulic properties for freshwater conditions. The ground-water and pure-water density and viscosity arrays are the output from a procedure (program STEAM) described on p. 18. The model grid spacing is that of a conventional discretization of the aquifer system for a horizontal plane at the surface.

In table 1, the DELTA array is a set of numerical factors for each node that may be used to modify model conductivities to account for possible decrease in conductivity as a function of overburden thickness. A discussion of this procedure, as used in the northern Midwest RASA simulation, is given in Mandle and Kontis (in press). As noted in table 1 and in program VARDEN (appendix B), if the constant IDELTA is set to zero and the DELTA arrays are set to unity, the conductivities will not be modified by the program.

The program output consists of the arrays given in table 1B and, if desired, a printed copy of all input and output data as well as the summation of source terms over each layer and over all layers. Information concerning the format of input and output arrays and details concerning the functioning of the program and its use are given in comment statements within the program listing (appendix B).

The modified Trescott-Larson program is based on equations with units of mass, length, and time in slugs, feet, and seconds, respectively. The procedure that computes density and viscosity (program STEAM) generates density in kilograms per cubic meter and viscosity in centipoise. In program VARDEN, ground-water density is read as kilograms per cubic meter and is converted first to grams per cubic centimeter and then to a relative density ( $\rho' = \text{DEN}$ ) by division by the ambient density ( $\rho_f$ ) of relatively fresh water in the study area as computed by program STEAM. For the northern Midwest RASA,  $\rho_f = 0.9977 \text{ g/cm}^3$ . Pure-water density ( $\rho_p$ ) is likewise converted to a relative density (DENP) when it is read and then is converted to density, in units of slugs per cubic foot. This is done by multiplying by the factor UNIT1, where:

$$\text{UNIT1} = 0.9977 \text{ g/cm}^3 \frac{(6.852 \times 10^{-5} \text{ slug/g})}{(3.531 \times 10^{-5} \text{ ft}^3/\text{cm}^3)} = 0.9977 \times 1.94 \text{ slug/ft}^3$$

Viscosity in centipoise is converted to a relative viscosity ( $\mu' = \text{VISC}$ ) by division by the ambient viscosity of relatively fresh ground water of the study area, as computed by program STEAM. The value for the northern Midwest RASA was 1.127 centipoise. If the model is to be executed in terms of freshwater head, the above quantities, as used in VARDEN, produce conductances in slugs per second per foot and source terms in slugs per second.

If the model is to be executed in terms of pressure (in slugs per foot second squared), the source terms of equations 2a-f are multiplied by a factor termed UNIT2, where:

$$\text{UNIT2} = \rho_f g = (1.94 \times .9977) \text{ slugs/ft}^3 \times 32.174 \text{ ft/s}^2,$$

and all conductances in equations 3a-f are divided by the same factor. Because pure-water density appears in the program as a relative density, the actual division is by  $g = 32.174 \text{ ft/s}^2$ . For the pressure case, VARDEN produces conductance in units of second-feet and source terms in units of slugs per second. The values of UNIT1 and UNIT2 factors are set in the program (appendix B).

To obtain the geometric terms required for equations 2 and 3, a local coordinate system  $\alpha, \beta, \gamma$  is developed at each node from specified horizontal grid spacings  $\Delta x_j$  and  $\Delta y_i$  in the  $x, y$  plane and elevations  $Z_{i,j,k}$  of the mid-points of the aquifer layers. At each node location in the  $x, y$  plane, the projection of the  $x$  axis, within the  $x, z$  plane, onto the plane of an aquifer layer defines the orientation of the  $\alpha$  axis. Thus, as shown in figure 2C, the  $\alpha$  axis is constrained to lie in the  $x, z$  plane. Similarly, the projection of the  $y$  axis onto the aquifer plane defines the  $\beta$  axis constrained to lie in the  $y, z$  plane. In figure 2C, let  $\vec{A}$  be a vector in the  $\alpha$  direction. The components of  $\vec{A}$  in the  $x$  and  $z$  directions are  $A_x$  and  $A_z$ , that is, for unit vectors  $\hat{i}$  and  $\hat{k}$  in the  $x$  and  $z$  directions.

$$\vec{A} = A_x \hat{i} + A_z \hat{k} = A \cos a' \hat{i} + A \sin a' \hat{k}.$$

The unit vector  $\hat{\alpha}$  in the  $\alpha$  direction is then:

$$\hat{\alpha} = \frac{\vec{A}}{|\vec{A}|} = \frac{A(\cos a' \hat{i} + \sin a' \hat{k})}{A(\cos^2 a' + \sin^2 a')^{1/2}} = \cos a' \hat{i} + \sin a' \hat{k}$$

In similar manner, a unit vector  $\hat{\beta}$  in the  $\beta$  direction (fig. 2B) in terms of unit vectors  $\hat{j}$  and  $\hat{k}$  in the  $y$  and  $z$  directions is given by

$$\hat{\beta} = \cos b' \hat{j} + \sin b' \hat{k}.$$

The vectors  $\hat{\alpha}$  and  $\hat{\beta}$  lie in the  $\alpha, \beta$  plane so that a vector  $\vec{\gamma}$  normal to the  $\alpha, \beta$  plane (fig. 2A) is the vector product of  $\hat{\alpha}$  and  $\hat{\beta}$ , or

$$\vec{\gamma} = \hat{\alpha} \times \hat{\beta} = -\cos b' \sin a' \hat{i} - \cos a' \sin b' \hat{j} + \cos a' \cos b' \hat{k}.$$

If  $\hat{\gamma}$  is a unit vector in the  $\gamma$  direction, and if  $c'$  is the angle between the  $\gamma$  and  $z$  axes then for small angles  $a', b'$ :

$$\cos c' = \hat{\gamma} \cdot \hat{k} = \frac{\vec{\gamma} \cdot \hat{k}}{|\vec{\gamma}|} = \frac{\cos a' \cos b'}{(1 - \sin^2 a' - \sin^2 b')^{1/2}} \approx \cos a' \cos b' = \frac{\Delta x \Delta y}{\Delta \alpha \Delta \beta}$$

This relation is used to approximate the gradient of elevation in the  $\gamma$  direction. As pointed out by Kuiper (1983), the  $\alpha$  and  $\beta$  directions defined in this manner are not, in general, normal to each other. For the small dips (generally less than several degrees) that usually occur in nature, the angle between  $\alpha$  and  $\beta$  is almost  $90^\circ$ , and the  $\alpha, \beta, \gamma$  system is essentially orthogonal.

For aquifers that dip uniformly, the geometric relations given below are exact, whereas for aquifers with small dip and where changes in dip are gradual the relations are a good approximation. Conversely, for strata characterized by large and rapidly changing dips, the computations of inclined block widths may be erroneous. With regard to the general problem of correctly simulating complex aquifer geometry, regardless of the density distribution, Weiss (1985)

developed the ground-water-flow equation for uniform-density conditions in curvilinear coordinates that follow the bedding of strata and that rigorously take into account changes in dip. By comparing the simulated flow system obtained from finite-difference approximations based on conventional x,y,z cartesian coordinates and on curvilinear coordinates for several test cases, Weiss concluded that the hydraulic effects of variable dip was relatively small, especially for dips less than several degrees.

The geometric terms consisting of inclined distances and inclined block widths  $\Delta\alpha$  and  $\Delta\beta$  are approximated in the program (appendix A) as follows. Figure 2C shows three adjacent nodes (i,j-1,k), (i,j,k) and (i,j+1,k) for a layer dipping in the  $\alpha$  direction with horizontal node spacing of  $\Delta x_{j-1}$ ,  $\Delta x_j$ , and  $\Delta x_{j+1}$ , respectively. The variables used in the program (appendix A) consist of the terms listed below.

DJM =  $(\Delta x_{j-1})/2$ ,  
 DJ =  $(\Delta x_j)/2$ ,  
 DJP =  $(\Delta x_{j+1})/2$ ,  
 DAM = inclined distance between node i,j-1,k and node i,j,k  
        $(\alpha_{i,j,k} - \alpha_{i,j-1,k})$ ,  
 DAP = inclined distance between node i,j,k and node i,j+1,k  
        $(\alpha_{i,j+1,k} - \alpha_{i,j,k})$ ,  
 DA = width of block i,j,k in  $\alpha$  direction  $(\Delta\alpha_{i,j,k})$ ,  
 DZAM = difference in elevation of nodes i,j-1,k and i,j,k,  
 DZAP = difference in elevation of nodes i,j,k and i,j+1,k.

From figure 2,

$$\begin{aligned} \text{DAM} &= ((\text{DJM} + \text{DJ})^2 + \text{DZAM}^2)^{1/2} \\ \text{DAP} &= ((\text{DJ} + \text{DJP})^2 + \text{DZAP}^2)^{1/2} \end{aligned}$$

If  $\theta$  is the angle between the vertical (z) and  $\alpha$  directions (fig. 2C),

$$\sin\theta = \text{DJ}/(\text{DA}/2) = (\text{DJM} + \text{DJ})/\text{DAM},$$

so that the block width is

$$\text{DA} = 2 * \text{DJ} * \text{DAM} / (\text{DJ} + \text{DJM}).$$

The block width for node i,j+1,k is

$$\text{DA2} = 2 * \text{DJP} * \text{DAP} / (\text{DJ} + \text{DJP}).$$

The block width  $\text{DB} = \Delta\beta$ , in the  $\beta$  direction, is developed in a similar manner from the terms of figure 2B. In the northern Midwest, the maximum dip of strata over model-node distances is less than 1.5 degrees. In general, for most aquifer systems, maximum dip of strata is probably less than several degrees, and the angles a',b',c' (fig. 2A) are quite small, so that slant distances between nodes and inclined block areas ( $\Delta\alpha$  and  $\Delta\beta$ ) and, in turn, model conductances TR, TC, and TK, will be close in magnitude to their corresponding values for an

x,y,z coordinate system. Consequently, elevation changes are significant only with respect to computation of the variable density terms A and B and less so for the  $\Gamma$  component (eq. 2f).

#### Description of Program STEAM

Ground-water density, pure-water density, and viscosity arrays for use in VARDEN may be obtained from a modified version (program STEAM) of a method developed and documented in detail by Weiss (1982a). The method is based on the assumption that the density of ground water is equivalent to the density of a sodium chloride solution of the same molality, temperature, and pressure. Under this assumption, the density of ground water at node i,j,k as a function of temperature at constant composition and pressure, can be computed through a relation from Potter and Brown (1977):

$$\rho_{s,i} = [C_a + C_b T_g + C_c T_g^2]_{i,j,k}, \quad (8)$$

where:

$T_g$  = temperature of ground water, in degrees Celsius, and

$C_a$ ,  $C_b$ ,  $C_c$  = empirical coefficients for sodium chloride solutions given in table 29 of Potter and Brown (1977).

The coefficients are given for values of molality ranging from 0.5 to 6.0 moles per kilogram (mol/kg) in increments of 0.5 mol/kg and of pressure ranging from standard atmospheric to 2,000 bars (29,000 lb/in<sup>2</sup>) in increments of 100 bars (1,450 lb/in<sup>2</sup>) up to 1,000 bars and 250 bars thereafter. Part of table 29 from Potter and Brown (1977) is reproduced in table 2. In addition, table 2 contains  $C_a$ ,  $C_b$ , and  $C_c$  coefficients for zero molality as determined by R. J. Sun (U.S. Geological Survey, written commun., 1981).

The molality for node i,j,k is computed iteratively by the relation (Weiss, 1982a).

$$XMOL_{i,j,k} = \frac{DS}{\left[ wt \left( DEN1 - \frac{DS}{1000} \right) \right]_{i,j,k}}, \quad (9)$$

where:

$XMOL$  = approximate molality of ground water, in mol/kg,

$DS$  = dissolved-solids concentration of ground water, in g/L at laboratory conditions (20 °C and 1 bar),

$wt$  = gram molecular weight of sodium chloride (58.4428 g/mol), and

$DEN1$  = density of ground water at 20°C and 1 bar, in g/cm<sup>3</sup>.



Table 2.--NaCl  $C_a$ ,  $C_b$ ,  $C_c$  interpolation coefficients

[From table 29 of Potter and Brown, 1977 and R. J. Sun, U.S. Geological Survey, written commun., 1981.  
SAT = standard atmospheric pressure.]

A. $C_a$ INTERPOLATION COEFFICIENTS						
I	Molality	Pressure (bars)				
		SAT J=1	100 J=2	200 J=3	300 J=4	400 J=5
1	0.0	1.00184	1.00667	1.01121	1.01548	1.01948
2	0.5	1.00590	1.02100	1.02600	1.03100	1.03400
3	1.0	1.03120	1.04300	1.04700	1.05100	1.05500
4	1.5	1.04900	1.06300	1.06800	1.07100	1.07500
5	2.0	1.07700	1.08100	1.08500	1.08800	1.09300
6	2.5	1.09900	1.09800	1.10200	1.10700	1.10900
7	3.0	1.11700	1.11600	1.11900	1.12300	1.12500
8	3.5	1.13700	1.13500	1.13500	1.13800	1.14000
9	4.0	1.15500	1.15000	1.15200	1.15400	1.15600
10	4.5	1.16800	1.16600	1.16800	1.17000	1.17200
11	5.0	1.18400	1.18300	1.18300	1.18500	1.18700
12	5.5	1.19800	1.19400	1.19700	1.19900	1.20100
13	6.0	1.21300	1.21000	1.21200	1.21500	1.21700

B. $C_b$ INTERPOLATION COEFFICIENTS TIMES -0.0001						
I	Molality	J=1	J=2	J=3	J=4	J=5
1	0.0	1.50194	1.29778	1.48770	1.65600	1.80564
2	0.5	-.31800	.23900	.86800	1.07700	.30500
3	1.0	.74300	1.20000	1.39900	1.41700	.84300
4	1.5	1.02800	1.61300	1.86900	1.87800	1.12100
5	2.0	2.76900	1.99400	1.60200	2.22200	1.41800
6	2.5	3.60200	1.66600	1.93900	2.63100	1.62200
7	3.0	4.13600	2.54800	2.23900	2.89400	1.79800
8	3.5	4.74600	3.43900	2.60800	2.85500	1.91400
9	4.0	5.10600	3.74100	3.35400	3.16100	2.23300
10	4.5	5.08400	4.10300	3.88100	3.68800	2.64400
11	5.0	5.33400	5.58800	3.96700	3.77200	2.79300
12	5.5	5.29200	4.26600	4.06600	3.87000	2.93200
13	6.0	5.42800	4.64200	4.46700	4.29400	3.54300

C. $C_c$ INTERPOLATION COEFFICIENTS TIMES -0.000001						
I	Molality	J=1	J=2	J=3	J=4	J=5
1	0.0	2.73006	3.14091	2.97626	2.83139	2.70357
2	0.5	3.09900	2.97400	2.59200	2.42700	2.56900
3	1.0	2.66100	2.52000	2.33600	2.21300	2.32600
4	1.5	2.41300	2.31400	2.10200	2.00100	2.43000
5	2.0	1.88800	2.27600	2.36800	1.80300	1.96700
6	2.5	1.60100	2.41100	2.17300	1.63500	1.84700
7	3.0	1.37800	1.95500	1.99400	1.49600	1.75200
8	3.5	1.17400	1.59300	1.82000	1.46400	1.60800
9	4.0	1.03300	1.31800	1.34300	1.33800	1.52500
10	4.5	.98000	1.16500	1.13700	1.13800	1.36900
11	5.0	.88700	.71700	1.09100	1.09000	1.30300
12	5.5	.85800	1.02800	1.03800	1.04300	1.24200
13	6.0	.80500	.88500	.89200	.89100	1.04700

In the program STEAM (appendix C) the iteration process is implemented by using an initial value of unity for DEN1 and computing molality from equation 9. The resultant value of molality is used in a relation applicable for constant temperature and pressure and for temperature less than 100 °C (eq. 1 of Potter and Brown, 1977). That is:

$$DEN1_{i,j,k} = \frac{(1000 + wt \cdot XMOL)DEN0}{1000 + (A_0 \cdot XMOL + B_0 \cdot XMOL^{3/2} + C_0 \cdot XMOL^2)DEN0}_{i,j,k}}, \quad (10)$$

where:

DEN0 = density of ground water at 20 °C, 1 bar

A<sub>0</sub>, B<sub>0</sub>, C<sub>0</sub> = empirical constants (see table 28, Potter and Brown, 1977).

For the next iteration, the value of density computed from equation 10 is used in equation 9 to recompute molality. This process is repeated until the percentage of molality difference between successive iterations is less than some specified closure criteria (CDIFF of Program STEAM).

For the northern Midwest study, estimates of the total dissolved-solids concentration were obtained from regional data compiled by D. I. Siegel (U.S. Geological Survey, written commun., 1981).

The pressure in bars at node i,j,k is assumed to be hydrostatic and computed from:

$$P_{i,j,k} = 980.665 \times 10^{-6} \text{ g/cm}^2\text{s}^2 \times 30.48 \text{ cm/ft} \times Z'_{i,j,k} \quad (11)$$

where:

Z'\_{i,j,k} = distance in feet from the surface to the midpoint of node i,j,k

The procedure used by Weiss (1982a) to obtain the C<sub>a</sub>, C<sub>b</sub>, and C<sub>c</sub> coefficients required for equation 8 is based on linear interpolation of table values for the computed molality and pressure values. An alternative method is used in the program STEAM. If each set of table values (tables 2A, 2B, and 2C) are considered to be a group of uniformly spaced discrete points from a continuous two-dimensional surface, then the C<sub>a</sub>, C<sub>b</sub>, and C<sub>c</sub> values for use in equation 8, which correspond to computed values of molality and pressure, may be obtained by bicubic interpolation. The procedure consists of fitting a bicubic polynomial (Davis and Kontis, 1970) for each cell defined by the four points (I,J), (I+1,J), (I,J+1), and (I+1, J+1), where I = 1,...,12, and J = 1,...,4 of tables 2A, 2B, and 2C. This results in a set of 16 bicubic coefficients for each cell of each table. The interpolation is effected by determining molality and pressure according to equations 9 and 11 and solving a bicubic polynomial. For example, the C<sub>a</sub> value for a particular value of pressure (x) and molality (y) is determined from the relation:

$$C_a(x,y)_{j,i} = \sum_{m=0}^3 \sum_{n=0}^3 BC_{m,n}^{j,i} \frac{(x-x_j)^m}{100} \frac{(y-y_i)^n}{0.5}, \quad (12)$$

where:

$x_j$  = tabulated pressure for column J (table 2A)

$y_i$  = tabulated molality for row I (table 2A) and

$BC_{mn}$  = 16 bicubic coefficients associated with cell J, I (table 2A).

For the same molality and pressure,  $C_b$  and  $C_c$  are obtained in the same manner from their respective sets of bicubic coefficients. The bicubic coefficients are not given in this report but are available upon request.

After determination of the  $C_a$ ,  $C_b$ , and  $C_c$  values, ground-water density is obtained from equation 8. For the northern Midwest, the temperature values  $T_g$  for use in equation 8 were obtained from a straight-line fit of temperature versus depth data for the Illinois basin (Bond, 1972) given by:

$$T_{g,i,j,k} = 0.0083Z'_{i,j,k} + 17.09 \text{ } ^\circ\text{C}$$

Given the density of ground water,  $\rho_g$ , the pure-water component is then obtained from:

$$\rho_{p,i,j,k} = \rho_{s,i,j,k} \left( 1 - \frac{1}{1 + \frac{1000}{XMOL \cdot wt}_{i,j,k}} \right) \quad (13)$$

The dynamic viscosity required in the variable-density formulation (eq. 3) is obtained from an empirical relation developed by Weiss (1982b) based on curves of viscosity versus temperature for various values of dissolved-solids concentration (Matthews and Russell, 1967). The relation is:

$$\mu_{i,j,k} = (38.3432/T_g^{1/2} - 14.621/T_g^{1/4} + 1.481)(1 + DS/300)|_{i,j,k}, \quad (14)$$

where:

$\mu_{i,j,k}$  = dynamic viscosity at node i,j,k, in centipoise, (M/L)/T,

$T_g$  = temperature, in  $^\circ\text{F}$ , and

DS = dissolved-solids concentration, as previously defined.

The procedure to compute equations 8-14 (program STEAM) is given in appendix C. Comment statements within the program explain the general functions of the program. Table 3 summarizes the required input data and the type of output generated by the program.

Table 3.--Program STEAM input and output arrays

CONTROL RECORD IN (F10.6,3I10) FORMAT		
Name in program	Description	Fortran unit number
CDIFF	Molality iteration closure criterion	5
IPRINT	$\left\{ \begin{array}{l} 1 \text{ if print of computations at each iteration} \\ \text{is desired.} \\ 0 \text{ otherwise} \end{array} \right.$	
IO	Number of model rows	
JO	Number of model columns	

## A. INPUT ARRAYS

Name in program	Description	Fortran unit number	Format	Units
TDS	Dissolved solids concentration	10	10F8.3	g/l
H	Freshwater heads	11	10F8.0	ft
ELEV	Elevation of nodes	12	10F8.0	ft
BC	Bicubic coefficients for $C_a$ table values	13	10F8.4	--
BC	Bicubic coefficients for $C_b$ table values	14	10F8.4	--
BC	Bicubic coefficients for $C_c$ table values	15	10F8.4	--

## B. OUTPUT ARRAYS

Name in program	Description	Fortran unit number	Format	Units
ZMU	Viscosity array	20	10F8.3	cp
XDEN	Groundwater-density array	21	10F8.1	kg/m <sup>3</sup>
XDENP	Pure-water-density array	22	10F8.1	kg/m <sup>3</sup>
Program printout		6		

## Water-Density Modifications to Trescott-Larson Program

Modifications of the Trescott-Larson model to implement equation 5a consist of insertion of statements to (1) directly read the TR and TC conductances computed by VARDEN, (2) bypass Entry TCOF in subroutine COEFF, wherein the values of TR and TC were formerly computed, and (3) increase the two-dimensional recharge array, QRE, to three dimensions to provide space for the variable-density source terms,  $TS_{i,j,k}$ . The specific computer-code modifications are given in appendix D. The statement numbers of appendix D follow the statement-numbering system of the original Trescott-Larson documentation (Trescott, 1975). Modifications were made to the main program and to subroutines DATAI, CHECKI, COEFF, and SOLVE as follows:

Program MAIN.--Modifications given in appendix D consist of statements to label the new input arrays (TR and TC), expand the size of the QRE array by the number of model layers to provide space for the total source arrays (TS), and change the call to Entry ARRAY of subroutine DATAI to read all input arrays from unit LIN. To read the TR and TC conductances, two additional calls to Entry ARRAY are performed.

Subroutine DATAI.--Modifications given in appendix D consist of statements to expand the QRE array to three dimensions and to read input arrays from unit LIN in the correct order.

Subroutine CHECKI.--Modifications given in appendix D consist of statements to expand the QRE array to three dimensions.

Subroutine COEFF.--Modifications given in appendix D consist of statements to expand the QRE array to three dimensions and to bypass computation of TR and TC conductances.

Subroutine SOLVE.--Modifications given in appendix F consist of statements to expand the numbering of the QRE array to three dimensions. Variable water-density modifications of subroutine SOLVE are indicated by dollar signs in columns 73-75.

To execute the Trescott-Larson program modified according to appendix D, all documented procedures given in Trescott (1975) and Trescott and Larson (1976) apply, with the following exceptions.

- (1) To invoke the variable-density option, specify the option formerly used to apply areal recharge to the top layer (IQRE). If simulation of areal recharge is desired, the areal recharge can be converted to a mass rate and added to the total-source array for the top layer.
- (2) Place starting heads (or pressure) and the following arrays from program VARDEN in a single file (LIN = 11): modified storage coefficient, modified transmissivity, TK conductance, total source term, TR conductance, and TC conductance. The order of the arrays following the starting heads (or pressure) are as given in table 1B.
- (3) Because all conductances are precomputed, model node-spacing arrays (DELX and DELY) are set to unity.

With these procedures, the model generates either freshwater head or pressure (depending on what option is selected in VARDEN). Given the computed heads (or pressure), and the conductances (TR, TC, and TK) and individual sources terms (A, B, and I) from VARDEN, flow may be computed from the appropriate form of equations 6a-c.

## Effects of Multiaquifer Wells

### Finite-Difference Equation

Bennett and others (1982) outline a method for accounting, approximately, for the effects of multiaquifer wells in three-dimensional ground-water-flow simulation. Their analysis is based on the assumptions that the area represented by any one model node contains only one multiaquifer well, that the well is exactly at the nodal point, and that the node spacing is uniform and radial symmetry prevails in the head distribution surrounding the well in each aquifer.

A graphic representation of flow and head around a multiaquifer well, with discharge  $Q_{wt}$  and open to four layers, is shown in figure 6 (p. 26). In the figure,  $h_w$  is the water level in the multiaquifer well, and  $Q_{i,j,k}$  is the flow between layer  $k$  and the well. For each layer, the head  $h_{i,j,k}$  computed in a simulation for node  $i,j,k$  is assumed to prevail in the aquifer along a radius of  $r_{ak}$  around the well and thus around the node  $i,j,k$  itself.

Mandle and Kontis (in press) show how the effects of multiaquifer wells may be generalized for the case where a model node has numerous wells and where subsets of the wells may be open to differing combinations of aquifers. In their development, all wells within a finite-difference block area  $\Delta x_j \cdot \Delta y_i$  are assumed to be uniformly distributed and are segregated into categories that include those wells open to a unique combination of aquifers. Each well in a given category is assumed to have the same discharge, radius, and water level. The equation, in finite-difference form, to describe total freshwater mass flow,  $\rho_p Q_{i,j,k}^N$ , between node  $i,j,k$  and the wells of all categories (eq. 15) is in a form compatible with the variable-density formulation (eq. 5a). In equation 15:

$\rho_p$  = density of pure water, which can be set at 1.94 slugs/ft<sup>3</sup>  
if wells are producing relatively fresh water;

$T_{i,j,k}^*$  = modified transmissivity of equation 7;

$r_w$  = radius of the wells;

$h_{f,i,j,k}$  = freshwater head in node  $i,j,k$ ;

$u_m$  = number of wells in category  $m$ , ( $m = 1, 2, 3, \dots N$ ) for  $N$  categories;

$Q_{Tm}$  = total well discharge from all wells of category  $m$  in the block area  $\Delta x_j \cdot \Delta y_i$ ,

$k_{max}$  = total number of model aquifer layers;

$W(k,m) = k$  if a category  $m$  well is open to aquifer layer  $k$ , or  
 $0$  if a category  $m$  well is not open to aquifer layer  $k$ ; and

$\delta[W(k,m)] = 1$  if  $W(k,m) = k$  or  $0$  if  $W(k,m) = 0$ .

The  $\delta$  function provides that, in summations over layers  $n$  ( $n = 1, 2, \dots, k_{\max}$ ), only layers open to the well are included, and, in summations over categories  $m$  ( $m = 1, 2, \dots, N$ ), only those categories with wells open to layer  $k$  are included.

$$\begin{aligned} \rho_p Q_{i,j,k}^N = & \frac{2\pi T_{i,j,k}^* h_{f,i,k}}{\ln\left(\frac{r_{a_i}}{r_w}\right)} \sum_{m=1}^N u_m \delta[W(k,m)] - \frac{2\pi T_{i,j,k}^*}{\ln\left(\frac{r_{a_i}}{r_w}\right)} \sum_{m=1}^N \left\{ u_m \delta[W(k,m)] \frac{\sum_{n=1}^{k_{\max}} \left[ \frac{T_{i,j,n}^* h_{f,i,n}}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]}{\sum_{n=1}^{k_{\max}} \left[ \frac{T_{i,j,n}^*}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]} \right\} \\ & + \frac{T_{i,j,k}^*}{\ln\left(\frac{r_{a_i}}{r_w}\right)} \sum_{m=1}^N \left\{ \frac{\rho_p Q_{i,j,k} \delta[W(k,m)]}{\sum_{n=1}^{k_{\max}} \left[ \frac{T_{i,j,n}^*}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]} \right\}, \end{aligned} \quad (15)$$

In equation 15,  $r_{a_k}$  (effective radius) is the radius of a circle around a well along which the head computed for node  $i,j,k$  is assumed to prevail and is given by:

$$r_{a_k} = \frac{\sqrt{s^2 N_k}}{4.81} = \frac{s}{4.81} \sqrt{N_k}. \quad (16)$$

where:

$s$  = minimum of the finite-difference block dimensions  $\Delta x_j$ ,  $\Delta y_i$ , and

$N_k$  = number of wells in block area  $\Delta x_j \cdot \Delta y_i$ .

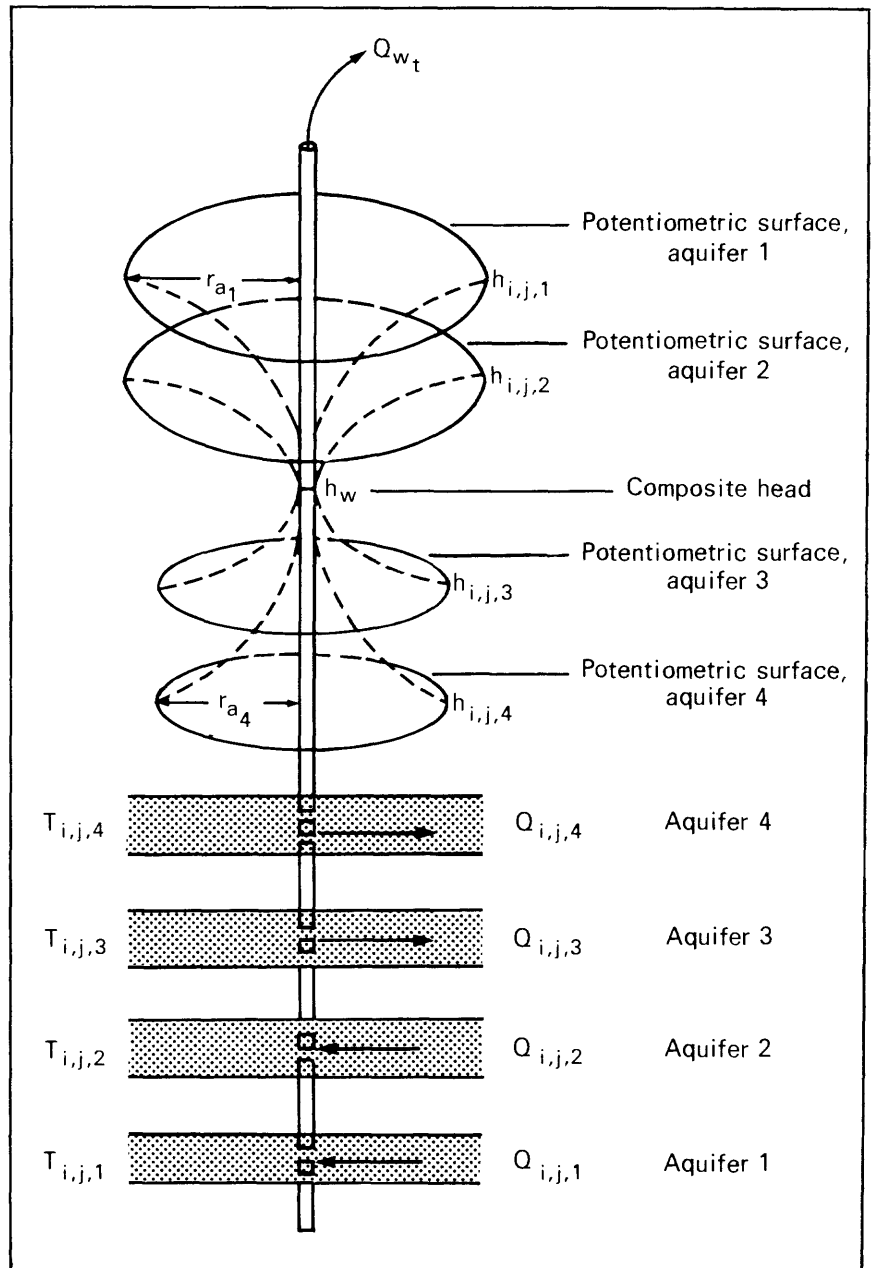
As discussed in Mandle and Kontis (in press) estimation of the radius,  $r_{a_k}$  (eq. 16), for each of the  $N_k$  wells is an oversimplification of actual field conditions. It is, however, a useful assumption when the purpose of a simulation is to approximate the effects of numerous multiaquifer wells on a flow system, as opposed to the detailed analysis of individual wells. The radius term is obtained by specifying that the sum of the assumed areas ( $N_k \pi r_{a_k}^2$ ) surrounding each of the  $N_k$  is equivalent to the area ( $\pi s^2 / 4.81^2$ ) for a single well of effective radius  $s/4.81$  (Prickett, 1967).

The water level in category m wells is given by:

$$h_{f_{n,m}} = \frac{\sum_{n=1}^{k_{mnx}} \left[ \frac{T_{i,j,n}^* h_{f_{i,j,n}}}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]}{\sum_{n=1}^{k_{mnx}} \left[ \frac{T_{i,j,n}^*}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]} - \frac{\rho_p Q_{T_n}}{2\pi u_m \sum_{n=1}^{k_{mnx}} \left[ \frac{T_{i,j,n}^*}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]}, \quad (17)$$

Figure 6.

Conditions of flow and head around a typical multiaquifer well.  
(From Bennett and others, 1982).





The generalized finite difference form of equation 5a, can be expressed as:

$$L^*(h_{f,i,j,k}) - \rho_p Q_{i,j,k}^N = 0,$$

where:  $L^*(h_{f,i,j,k})$  includes all terms of equation 5a except the product of density and flow to wells, which is now represented by the term  $\rho_p Q_{i,j,k}^N$  of equation 15.

The finite-difference formulation of equation 5a allows the pure-water density ( $\rho_p$ ) to be treated as a variable in space to accomodate areas of highly saline brine, where the concentration of dissolved salts may be so large that the mass of pure water per unit fluid volume is reduced. If most of the wells in a study area produce from zones of relatively fresh water, in which spatial variation of  $\rho_p$  is minimal, then a constant  $\rho_p$  value of 1.94 slugs/ft<sup>3</sup> may be used in all well-flow terms. For uniform-density applications, equation 15 applies directly if  $\rho_p$  is set to unity and  $T^*$  is set to the product of hydraulic conductivity and layer thickness.

A major advantage of using equation 15 in simulation of systems containing multiaquifer wells is that water will be extracted from each aquifer open to a well in proportion to its relative transmissivity. Otherwise, the contribution from each aquifer to total well discharge must be estimated and altered with each change in model transmissivity. Implementation of equation 15 eliminates the necessity of this step and allows rapid calculation of composite head with equation 17.

In equation 15, the first and second terms consist of various coefficients that multiply the heads of nodes that are linked to each other by the presence of multiaquifer wells. Because these links occur only at nodes above and below a given node, the slice-successive-over-relaxation (SSOR) method was chosen to solve the finite-difference equations to obtain the model heads. The SSOR method readily accommodates the multiaquifer terms because a solution coefficient matrix is constructed for vertical slices along model rows. The nonzero terms of the matrix represent conductances of the nodes within the slice. Consequently, to implement equation 15, the Trescott-Larson program was modified to compute the multiaquifer coefficients of the equation and to insert them in the proper location of the SSOR coefficient matrix. The last term in equation 15 is independent of head and, inserted into the well array of the Trescott-Larson program code, serves as a source/sink term in the finite-difference equation.

#### Modifications to Trescott-Larson Program

The correspondence between Fortran variables in the modified Trescott-Larson program and terms of the multiaquifer-well formulation (eq. 15) is given in table 4. The indices of the Fortran variables in the table are symbolic and represent: layer number  $K$ , ( $K = 1, 2, \dots, K_0$ ) for  $K_0$  layers; node-set number  $L$ , ( $L = 1, 2, \dots, N_{WEL}$ ) for  $N_{WEL}$  node sets (a node set consists of a vertical stack of nodes at a particular row and column intersection  $(I, J)$  for which a well or wells taps at least one aquifer); and well category number  $N$ , ( $N = 1, 2, \dots, J_A$ ) for  $J_A$  categories of a given node set.

Table 4.--Correspondence between Fortran variables used in  
Trescott-Larson model program and terms of  
multiaquifer-well formulation

Multiaquifer-well term (equation 15)	Fortran variable (equation 18)
kmax	K0
$W(k,m)$	KLL(K,L,N)
$Q_{T_m}$	QWL(N,L)
$r_{a_k}$	REL(K,L)
$r_w$	RWL(L)
$\rho_p$	1.94
$2\pi$	PI2
$T^*_{i,j,k}$	T(I,J,K)
$h_{f_{i,j,k}}$	PHI(I,J,K)
$u_m$	NU(N,L)
N	NWL(L)=JA
$1/\ln(r_{a_k}/r_w)$	C1(K,L)
$1 / \sum_{n=1}^{k_{max}} \left[ \frac{T^*_{i,j,n}}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]$	C2(N,L)
$\frac{T^*_{i,j,k}}{\ln\left(\frac{r_{a_k}}{r_w}\right)} \sum_{m=1}^N \left\{ \frac{\rho_p Q_{T_m} \delta[W(k,m)]}{\sum_{n=1}^{k_{max}} \left[ \frac{T^*_{i,j,n}}{\ln\left(\frac{r_{a_n}}{r_w}\right)} \right] \delta[W(n,m)]} \right\}$	WELL(I,J,K)

Through the Fortran variable notation given in table 4, the multiaquifer-well formulation (eq. 15) for node (I,J,K) corresponding to node set L may be written as:

$$\begin{aligned} \rho_p Q(I,J,K)^N = & \sum_{N=1}^{JA} NU(N,L) \cdot \delta[KLL(K,L,N)] \{ PI2 \cdot C1(K,L) \cdot T^*(I,J,K) \cdot PHI(I,J,K) \\ & - PI2 \cdot C1(K,L) \cdot T^*(I,J,K) \sum_{M=1}^{KO} PHI(I,J,M) \cdot T^*(I,J,M) \cdot C1(M,L) \cdot C2(N,L) \\ & \cdot \delta[KLL(M,L,N)] + WELL(I,J,K) \}. \end{aligned} \quad (18)$$

Equation 18 is the working formula for the multiaquifer-well terms that was incorporated in the Trescott model. The input data required by equation 18 are read in subroutine DATAI near the location where pumping rates were formerly read. A multiaquifer-well record contains the following information for each node set L:

IL(L) = row number,  
 JL(L) = column number,  
 NWL(L) = JA = total number of categories of multiaquifer wells,  
 KLL(K,L,N) = A KO digit code for each category N (N = 1,2...JA) of wells containing the layer number K if the layer is open to a well and zero otherwise, and  
 QWL(N,L) = total pumping rate, in million gallons per day, of all wells in category N.

For example, table 5 shows the input record for a node set L at model row 20 and column 27 that contains three categories of wells (NWL). The first category is open to layers 1, 2, 3, and 4; the second to layers 1 and 2; and the third to layer 1 only. The pumping rates for each category are 2, 1 and 0.5 million gallon per day, respectively. The multiaquifer-well input file consists of a record similar to the table below for each node set and is ordered by column within each row starting with the topmost row.

Table 5.--Example of multiaquifer-well input record  
 for node set L at row 20, column 27.

INPUT-VARIABLE NAME								
IL(L)	JL(L)	NWL(L)	KLL(K,L,1)	KLL(K,L,2)	KLL(K,L,3)	QWL(1,L)	QWL(2,L)	QWL(3,L)
20	27	3	1234	1200	1000	2.0	1.0	0.5

The multiaquifer well formulation was incorporated in the Trescott-Larson model as an additional option so that the program can be executed in the conventional manner if the multiaquifer-well option is not desired. The following routines of the program were changed:

- (1) MAIN program and subroutine PRNTA to include the new multiaquifer-well option,

- (2) subroutine DATAI to read the multiaquifer-well input data (table 5) and to compute some of the terms of table 4,
- (3) subroutine CHECKI to compute and output composite heads, and
- (4) subroutine SOLVE (SSOR solution method) to call two new subroutines, MULT and RESI.

The new subroutine MULT computes modified and additional coefficients resulting from the presence of multiaquifer wells and places them in the proper location of the SSOR solution coefficient matrix. Subroutine RESI computes the multiaquifer-well contributions for the "right-hand" side of the solution matrix equation. In addition, a third new subroutine ALOCAT computes all flows into and out of each node of a multiaquifer node set and a balance for each node.

A brief discussion of multiaquifer-well modifications of the Trescott-Larson routines follows. A discussion of the SSOR method (subroutines SOLVE and BAND), and a discussion of the new subroutines (MULT, RESI, and ALOCAT) is given in appendix A; and a listing of these routines is given in appendix F and G. A listing of the multiaquifer-well modifications is given in appendix E; the statement numbers of appendix E follow the numbering system of the original Trescott documentation (Trescott, 1975).

Program MAIN.--Modifications, given in appendix E, consist of adding a common block termed WIS containing the multiaquifer-well input variables of table 4, dimensioning new variables with a parameter statement, adding two new options (IOPEN and IALOCA) to the model options list, and specifying the value of PI2. In addition, several format changes and changes to the call to entry ARRAY of subroutine DATAI to specify a logical unit number (LIN) for input data were made. To invoke the multiaquifer-well option and subroutine ALOCAT, the words "OPEN" and "ALOC" are specified in the 12th and 13th fields of the Trescott-Larson options record.

Subroutine DATAI.--Modifications, given in appendix E, consist of including the new options (IOPEN and IALOCA), common-block WIS, the logical unit number LIN, code to read the multiaquifer-well records (as in table 5), and statements to compute or specify the terms REL, RWL, NU, C1, C2 and WELL of table 4. Because the northern Midwest pumping-rate file contains positive values for discharge (table 5), the pumping rate input (QWL), in million gallons per day, is multiplied by -1 and converted to cubic feet per second within the subroutine. If a simulation includes the variable water-density option, the pumping rates are multiplied by the density of pure water (1.94 slugs/ft<sup>3</sup>) to convert to units of mass/time required by the variable water-density formulation. The number of wells (NU) in a given category was estimated by dividing the pumping rate (QWL) by an assumed pumping rate per well of 0.1 ft<sup>3</sup>/s, and the well radius (RWL) was taken to be 0.5 ft.

The effective radius ( $r_{ak}$  or REL of table 4) of wells tapping layer k is estimated in the subroutine as follows:

From equation 16,

$$r_{ak} = \frac{s}{4.81} \sqrt{\frac{1}{N_k}}$$

$N_k$ , the total number of wells open to layer K for a particular node set L, is computed by

$$N_k = NR(K,L) = \sum_{JJ=1}^{JA} NU(JJ,L) \cdot \delta[KLL(K,L,JJ)],$$

where NU and JA are as previously defined (table 4).  
Consequently,

$$r_{ak} = REL(K,L) = \sqrt{\frac{s^2}{4.81^2 NR(K,L)}} = \sqrt{\frac{0.04322 s^2}{NR(K,L)}}.$$

The multiaquifer-well coefficients are not particularly sensitive to variations in the value of  $r_{ak}$  and  $r_w$  because the effective and well radii appear in equation 15 as the natural logarithm of their ratio ( $r_{ak}/r_w$ ).

Subroutine CHECKI.--Modifications, given in appendix E, were made to include the new options (IOPEN and IALOCA) and common block WIS, dimension arrays HW and HZ used to compute composite head; output the composite head and call subroutine ALOCAT if the IALOCA option is invoked. In subroutine CHECKI, HW is the composite head given by equation 17, whereas HZ is the composite head with no pumping ( $Q_{T_m} = 0$ ) and represents model results that may be compared with measured data from observation wells. For the latter case, the output consists of records similar in structure to the multiaquifer-well input records (table 5). For each node set the record contains the row and column, number of multiaquifer-well categories, the layer code, and the observation-well composite-head for each category. These records are written to Fortran unit 13.

Subroutine PRNTA.--Modifications, given in appendix E, consist of adding the new options IOPEN and IALOCA.

## SUMMARY

The ground-water flow equation used in the Northern Midwest Regional Aquifer-System Analysis was modified to simulate the effects of time invariant, spatially varying ground-water density, and wells open to more than a single aquifer. Corresponding modifications of the Trescott-Larson three-dimensional ground-water flow program to allow simulation of these effects are presented. In addition, the functioning of the slice-successive-overrelaxation method to solve for head distribution is summarized and illustrated. Two new programs are presented: program STEAM computes ground-water density, pure-water density, and dynamic viscosity; program VARDEN computes model-conductance and variable-density input data for the Trescott-Larson model.

The following steps are required to simulate variable-density effects with the modified Trescott-Larson model (appendix D):

- (1) For each model layer, execute program STEAM (appendix C) with the input data given in table 3 to generate arrays of ground-water density, pure-water density, and dynamic viscosity.
- (2) From the output of program STEAM (table 3) and the additional input given in table 1, execute program VARDEN (appendix B) to generate input data for the Trescott-Larson model (table 1) for the terms of equation 5a and individual variable-density terms (table 1) required to compute flow (eq. 6).
- (3) Place model starting heads and output arrays of program VARDEN (table 1) in a single file (Fortran unit-number LIN of appendix D).
- (4) Set model grid spacing (DELX and DELY) to unity, specify the IQRE option of Trescott-Larson model and execute the model in the conventional manner (as given in Trescott, 1975) to generate freshwater head or pressure.

The following procedures are required to simulate the effects of multi-aquifer wells with the modified Trescott-Larson (appendix E) and SSOR subroutines (appendix F), and the new subroutines MULT and RESI (appendix G).

- (1) If variable density is not being simulated or if model grid spacing is nonuniform modify statements DAT1790 and Dat1797 according to comments in subroutine DATAI (statements DAT1786 and DAT1793 of appendix E) pertaining to multiplication of pumping rates by density of pure water and computation of area of block (AREA).
- (2) Create the multiaquifer pumping-rate file to incorporate records similar to those in table 5 for each model node set containing a well(s). Records are ordered columnwise starting at top row.
- (3) Specify an output file (Fortran unit 13) to save resultant multiaquifer composite heads (eq. 17).

- (4) Invoke IOPEN option in the 12th field of the Trescott-Larson options record and execute the Trescott-Larson model in the conventional manner (as given in Trescott, 1975).

Files of programs discussed in this report are available from the authors upon request.

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## APPENDIX A.--SLICE-SUCCESSIVE-OVER-RELAXATION (SSOR) SOLUTION METHOD

A narrative of the major elements of the SSOR procedure used in this study follows, and a listing of subroutine SOLVE, which implements the SSOR method, and subroutine BAND, which solves the set of resultant linear equations by Gaussian elimination, is given in appendix F. The description of the SSOR method is presented to illustrate how the Trescott-Larson model was modified to incorporate the effects of multiaquifer wells. In subroutine SOLVE (appendix F), original program statements as written by S. P. Larson and C. R. Faust (U.S. Geological Survey written commun., 1979), multiaquifer-well statements, and comments describing the functioning of the subroutine are indicated by "SP3," asterisks, and zeros in columns 73-75, respectively. Subroutine SOLVE contains two entry points--entry ITER and entry NEWITA. Entry ITER, entered once per simulation, contains statements to (1) identify active model nodes, (2) number nodes, and (3) order and count the solution equations for each slice and for the entire model. Entry NEWITA, entered each time step, sets up matrices required to obtain solutions for head and calls subroutine BAND to implement the solution.

### Entry ITER

In the SSOR routine, model nodes are numbered according to a one-dimensional numbering convention. The convention is illustrated for a single slice through a cross-sectional model consisting of 3 rows, 5 columns, and 5 layers in figure A-1. The node numbering begins at the upper left corner of the bottom layer ( $K = 1$ ) and proceeds in sequence along model rows. In Entry ITER, given in appendix F, the node number is designated by the variable N. Of the 75 model nodes in the example (fig. A-1), 15 are active (circled) because the border nodes of each layer are inactive. The number of active nodes and, therefore, the number of equations for each slice are counted and placed in array NB. For the example (fig. A-1), NB(I) for  $I = 2$  is 15. The equations for each slice are ordered in array NP by indexing on layer (K) and then on column (J). If a node is inactive the corresponding value of NP is zero. Table A-1 illustrates the equation ordering system for the 15 active nodes of the slice through  $I = 2$ . In addition to numbering nodes and ordering and counting equations for each slice, the maximum half-band width (MBW) of the solution coefficient matrix is computed (MBW = number of layers + 1). The SSOR acceleration parameter (WO), usually chosen to be some value between 1 and 2, and the number of iterations between application of the so-called Watt's correction (LENGTH), are read in Entry ITER.

### Entry NEWITA

For the conventional finite-difference computation stencil shown in figure A-2A (p.38), the finite-difference equation of node  $i,j,k$  at the start of a time step and with no multiaquifer well is:

$$\begin{aligned} & Eh_{i,j,k} + Dh_{i,j-1,k} + Fh_{i,j+1,k} + Bh_{i-1,j,k} + Hh_{i+1,j,k} \\ & + Zh_{i,j,k-1} + SUh_{i,j,k+1} = \frac{RHO}{\Delta t} h_{i,j,k}^* + Q_{i,j,k}, \end{aligned} \quad (A-1)$$

Table A-1

Slice-successive-over-relaxation equation numbering scheme for cross-sectional model on right.

Node number (N)	Node (I,J,K)	Equation number (NP)
5	2,2,1	1
20	2,2,2	2
35	2,2,3	3
50	2,2,4	4
65	2,2,5	5
8	2,3,1	6
23	2,3,2	7
38	2,3,3	8
53	2,3,4	9
68	2,3,5	10
11	2,4,1	11
26	2,4,2	12
41	2,4,3	13
56	2,4,4	14
71	2,4,5	15

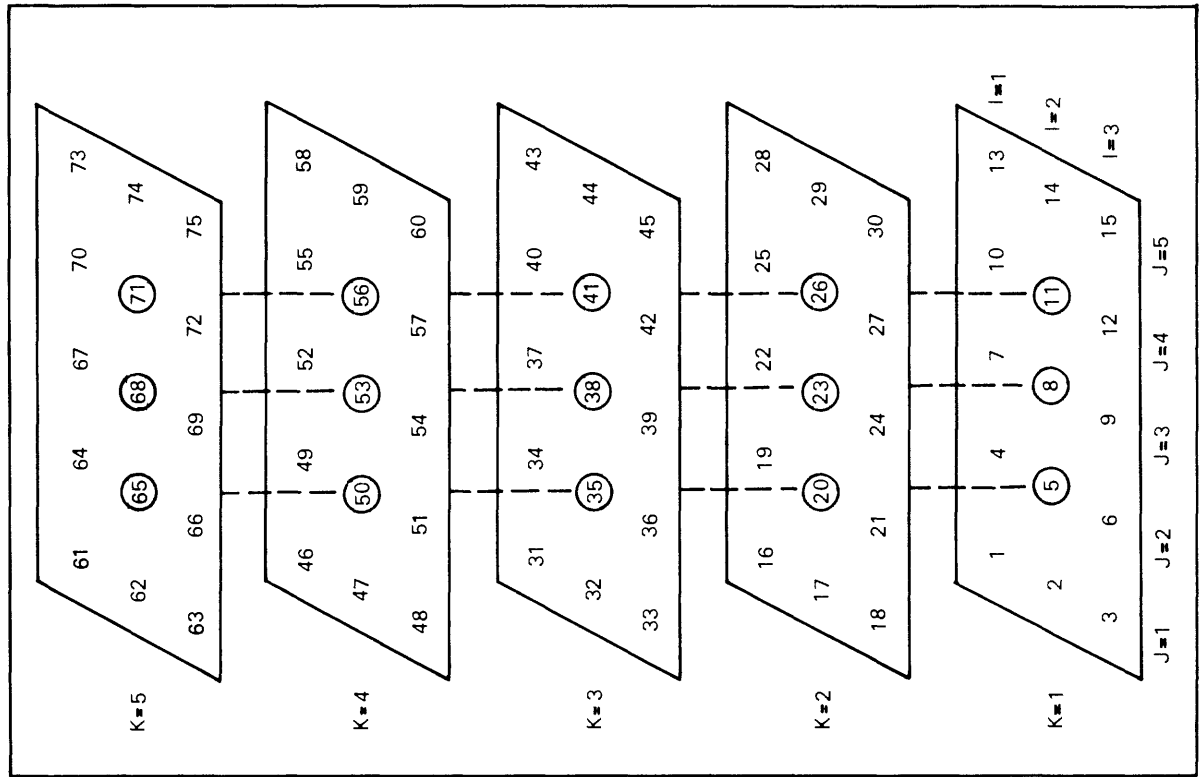


Figure A-1.

Example of slice-successive-over-relaxation node numbering scheme (values of N) for cross-sectional model consisting of 3 rows, 5 columns, and 5 layers.

where:

coefficients D, F, B, H, Z, and Su are the conductances between node i,j,k and the six adjacent nodes (fig. A-2),

$RHO/\Delta t$  is the storage coefficient times the area of the block i,j,k divided by the length of the time step,

Q represents sources and sinks, and

$h^*$  is the head at the beginning of the time step.

The coefficient of the head at i,j,k (E) is the sum of all conductances plus  $RHO/\Delta t$ .

In the SSOR routine, the solution is posed in terms of head differences rather than absolute head so that if  $\Delta h_{i,j,k}$  is defined to be the difference in head between time steps  $\tau+1$  and  $\tau$ , the finite-difference equation for node i,j,k at time step  $\tau+1$  may be written as

$$\begin{aligned} & E\Delta h_{i,j,k}^{\tau+1} + D\Delta h_{i,j-1,k}^{\tau+1} + F\Delta h_{i,j+1,k}^{\tau+1} + B\Delta h_{i-1,j,k}^{\tau+1} + H\Delta h_{i+1,j,k}^{\tau+1} + Z\Delta h_{i,j,k-1}^{\tau+1} + SU\Delta h_{i,j,k+1}^{\tau+1} \\ &= Q_{i,j,k} - Eh_{i,j,k}^{\tau} - Dh_{i,j-1,k}^{\tau} - Fh_{i,j+1,k}^{\tau} - Bh_{i-1,j,k}^{\tau} - Hh_{i+1,j,k}^{\tau} - Zh_{i,j,k-1}^{\tau} - SUh_{i,j,k+1}^{\tau} - \frac{RHO}{\Delta t}h_{i,j,k}^* \\ &= RHS_{i,j,k}^{\tau} \end{aligned} \quad (A-2)$$

The right hand side of equation A-2 consists of the product of coefficients and heads of the previous time step ( $\tau$ ) plus the source/sink terms and is called  $RHS_{i,j,k}$ . Equation A-2 is implemented according to the matrix relation

$$[A_s]^{\tau} [\Delta h]^{\tau+1} = [RHS]^{\tau}, \quad (A-3)$$

Where:

$[A_s]$  (array 'A') = symmetric compressed coefficient matrix of dimension  $NEQ \times MBW$  for  $NEQ$  active model nodes and half-bandwidth  $MBW$ ,

$[\Delta h]$  = vector of head differences of dimension  $NEQ$ , and

$[RHS]$  = vector of right-hand side terms of equation A-2 and is of dimension  $NEQ$ .

For the example (fig. A-1),  $MBW = 6$  and  $NEQ = 15$ .

In entry NEWITA, given in appendix F, the coefficients for the 'A' array and the RHS term are built according to the numbering scheme of table A-1 for all nodes containing a nonzero value in array NP. These terms are developed for each slice, once per time step, in statements SP30780 through SP31300 of subroutine SOLVE (appendix F).

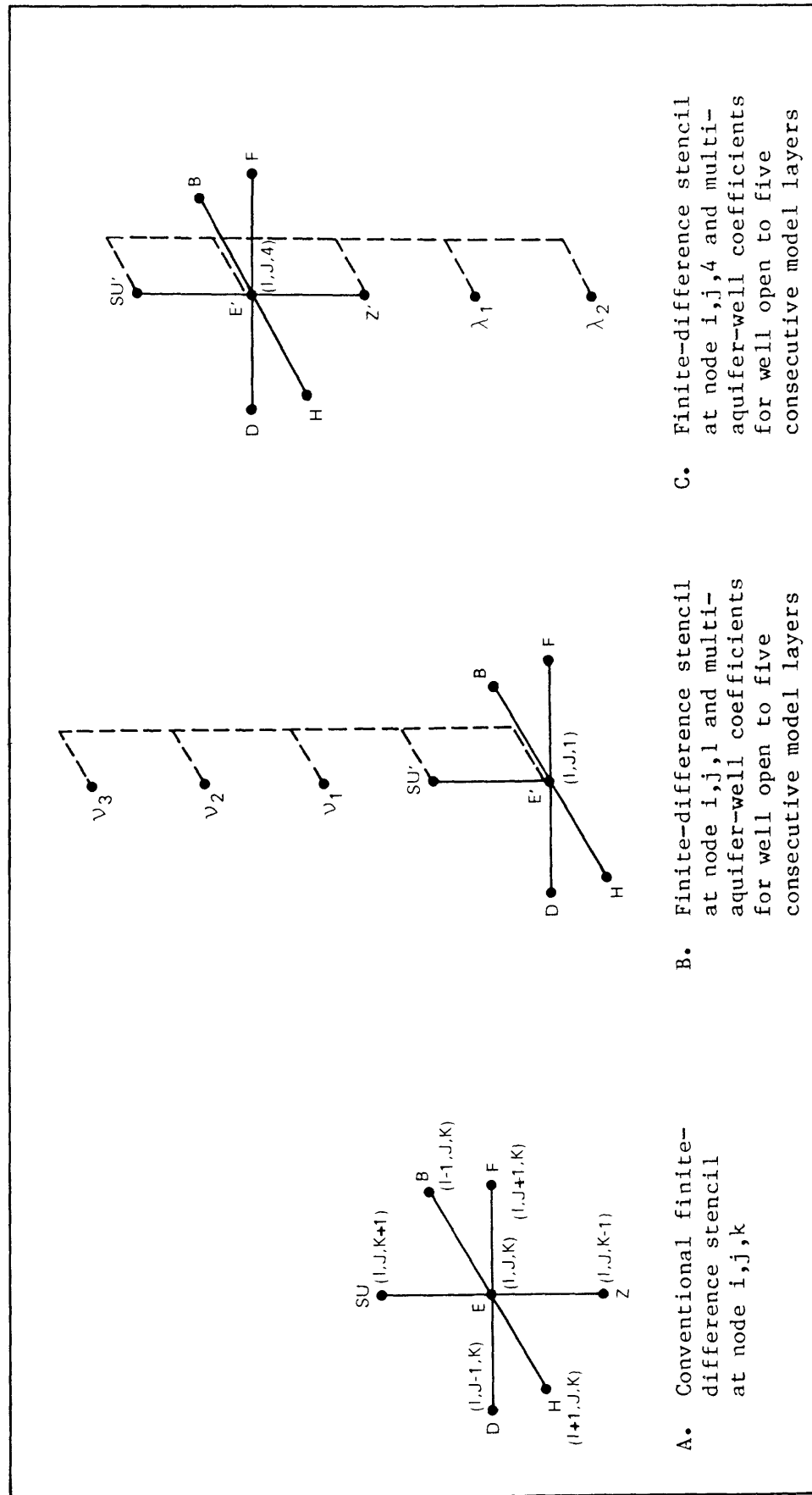


Figure A-2. Conventional finite-difference computation stencil and example of multi-aquifer-well stencil for five-layer case positioned at nodes in layers 1 and 4.

The solution for head differences at time step  $\tau+1$  is obtained by way of an iteration process. Because the SSOR procedure solves for the head differences within a slice, terms involving head differences for slice  $i-1$  ( $B\Delta h_{i-1,j,k}$ ) and slice  $i+1$  ( $H\Delta h_{i+1,j,k}$ ) (fig. A-2A) may be moved to the right side of the equation. Thus, for iteration  $m+1$  of time step  $\tau+1$  the equation for node  $i,j,k$  is:

$$E\Delta h_{i,j,k}^{m+1} + D\Delta h_{i,j-1,k}^{m+1} + F\Delta h_{i,j+1,k}^{m+1} + Z\Delta h_{i,j,k-1}^{m+1} + SU\Delta h_{i,j,k+1}^{m+1} =$$

$$RHS_{i,j,k}^{\tau} - (B\Delta h_{i-1,j,k}^{m+1} + H\Delta h_{i+1,j,k}^m) = RSMAL_{i,j,k}^{m+1}, \quad (A-4)$$

where  $RSMAL_{i,j,k}$  is the new right side of the equation incorporating the RHS term of equation A-2 and the coefficient terms involving B and H. Because the head differences for slice  $i-1$  are available at iteration  $m+1$ , they are used in the development of  $RSMAL$ . In matrix form, the solution equation for time step  $\tau+1$  and iteration  $m+1$  is

$$[A_i]^T [\Delta h]_{m+1}^{\tau+1} = [RSMAL]_{m+1} = [RHS]^T - [B\Delta h]_{m+1} - [H\Delta h]_m. \quad (A-5)$$

Application of equation A-4 and the nature of the compressed coefficient matrix  $A_S$  (array 'A') for the cross sectional slice through  $I = 2$  is illustrated in figure A-3A. In the figure, all elements having the same symbol will, in general, have different values. If the coefficient matrix were written in its full form, it would be of size  $NEQ \times NEQ$ , and the E coefficient associated with node  $i,j,k$  would lie along the main diagonal of the full matrix. Because the SU conductance of node  $i,j,k$  is equivalent to the Z conductance of node  $i,j,k+1$ , and the F conductance of node  $i,j,k$  is equivalent to the D conductance of node  $i,j+1,k$ , the full coefficient matrix is symmetric so that only coefficients for the upper diagonal part (E, SU and F) are required.

This symmetry allows development of the matrix in compressed form such that the main diagonal coefficients are located in the first column of each matrix row (fig. A-3A). The vector of head differences  $[\Delta h]$ , contains all the active nodes of the model and is ordered as to table A-1. Multiplication of components of  $[\Delta h]$  by the compressed coefficient matrix according to the relation

$$\sum_{i=1}^{MBW} A_{j,i} \Delta h_{j+i-1} = RSMAL_j, \quad j=1,2,...,NEQS$$

generates the set of equations to be solved for head difference for each iteration.

For each iteration, the  $RSMAL$  term (eq. A-4) is updated and subroutine BAND is entered to solve for an interim set of head differences, for each slice, by Gaussian elimination. The final head differences for iteration  $m+1$  are obtained by applying an over-relaxation factor (acceleration parameter  $W0$ ) to speed the rate of convergence--that is, for node  $i,j,k$

$$\Delta h_{i,j,k}^{m+1} = \Delta h_{i,j,k}^m + w0(\Delta h_{i,j,k}^{*m+1} - \Delta h_{i,j,k}^m), \quad (A-6)$$

where:

$\Delta h^*_{i,j,k}$  = interim head-difference solution, and

$\Delta h^m$  = solution from the previous iteration.

In subroutine SOLVE (appendix F), the iteration loop extends from statement SP31310 to SP32330. The section where the RSMAL terms are developed for each slice and where the interim solutions for head differences within the slice are obtained extends from statement SP31400 through SP31570. Application of the over-relaxation factor to obtain the head difference solutions extends from SP31580 to SP31690.

After iteration  $m + 1$  is completed for all slices, the maximum change in the head difference for all nodes, relative to the previous iteration, is compared with the specified closure criteria (ERR). If closure is achieved, final heads for the time step are formed by adding the head differences to the heads of the previous time step. This is performed from statements SP32350 to SP32400 (appendix F).

If closure is not achieved, a process termed the Watts correction, designed to speed the convergence, may be applied. In this process, a correction factor is computed for each slice, such that when the factor is added uniformly to the head-difference solution, the sum of the residuals over each slice is reduced to zero. If, however, a slice contains at least one constant-head node, the correction factor is zero, so that nothing is gained by applying the process to the slice. The Watts correction was not implemented for the northern Midwest study because the top layer consists of constant-head nodes.

#### Modifications of SSOR for Effects of Multiaquifer Wells

The multiaquifer-well modifications of the SSOR procedure in subroutine SOLVE consist of determining the set of coefficients as prescribed by equations 15 or 13 and placing them in their proper location in the solution coefficient matrix (array 'A') and the right-hand side vector RHS of equation A-2. No changes are required in the development of the RSMAL vector (eq. A-4) because the presence of multiaquifer wells does not alter the B and H coefficients.

The coefficients in question are those of equation 15 or 18 that multiply the heads of layers open to a well. For example, the multiaquifer-well stencil-position for a node set at layer 1 of a five-layer model where all layers are open to a well is shown in figure A-2B. The coefficients associated with layers 1 and 2 will consist of the E and SU terms of the standard stencil (fig. A-2A) plus a contribution from equation 15 or 18 forming coefficients E' and SU' (fig. A-2B). As shown in figure A-2B, new coefficients associated with layers 3, 4, and 5 ( $v_1$ ,  $v_2$ , and  $v_3$ ) are required. The set of coefficients for the computation stencil at a node in layer 4 is shown in figure A-2C. Because the multiaquifer-well coefficients are symmetric, the  $Z'$ ,  $\lambda_1$  and  $\lambda_2$  terms need not be determined explicitly. In the example, the  $Z'$ ,  $\lambda_1$  and  $\lambda_2$  coefficients of figure A-2C are equivalent to the SU',  $v_1$  and  $v_2$  coefficients when the com-

' A' ARRAY						$\Delta H(N), (I, J, K)$			RSMAL(N)		
-----						-----			-----		
E	SU	0	0	0	F	x	$\Delta H(5), (2, 2, 1)$		=	RSMAL(5)	
E	SU	0	0	0	F		$\Delta H(20), (2, 2, 2)$			RSMAL(20)	
E	SU	0	0	0	F		$\Delta H(35), (2, 2, 3)$			RSMAL(35)	
E	SU	0	0	0	F		$\Delta H(50), (2, 2, 4)$			RSMAL(50)	
E	0	0	0	0	F		$\Delta H(65), (2, 2, 5)$			RSMAL(65)	
E	SU	0	0	0	F		$\Delta H(8), (2, 3, 1)$			RSMAL(8)	
E	SU	0	0	0	F		$\Delta H(23), (2, 3, 2)$			RSMAL(23)	
E	SU	0	0	0	F		$\Delta H(38), (2, 3, 3)$			RSMAL(38)	
E	SU	0	0	0	F		$\Delta H(53), (2, 3, 4)$			RSMAL(53)	
E	0	0	0	0	F		$\Delta H(68), (2, 3, 5)$			RSMAL(68)	
E	SU	0	0	0	0		$\Delta H(11), (2, 4, 1)$			RSMAL(11)	
E	SU	0	0	0	0		$\Delta H(26), (2, 4, 2)$			RSMAL(26)	
E	SU	0	0	0	0		$\Delta H(41), (2, 4, 3)$			RSMAL(41)	
E	SU	0	0	0	0		$\Delta H(56), (2, 4, 4)$			RSMAL(56)	
E	0	0	0	0	0		$\Delta H(71), (2, 4, 5)$			RSMAL(71)	

(A) SSOR SOLUTION MATRICES -- NO MULTIAQUIFER WELLS.

' A' ARRAY						$\Delta H(N), (I, J, K)$			RSMAL(N)		
-----						-----			-----		
E'	SU'	$v_1$	$v_2$	$v_3$	F	x	$\Delta H(5), (2, 2, 1)$		=	RSMAL(5)	
E'	SU'	$v_1$	$v_2$	0	F		$\Delta H(20), (2, 2, 2)$			RSMAL(20)	
E'	SU'	$v_1$	0	0	F		$\Delta H(35), (2, 2, 3)$			RSMAL(35)	
E'	SU'	0	0	0	F		$\Delta H(50), (2, 2, 4)$			RSMAL(50)	
E'	0	0	0	0	F		$\Delta H(65), (2, 2, 5)$			RSMAL(65)	
E	SU	0	0	0	F		$\Delta H(8), (2, 3, 1)$			RSMAL(8)	
E	SU	0	0	0	F		$\Delta H(23), (2, 3, 2)$			RSMAL(23)	
E	SU	0	0	0	F		$\Delta H(38), (2, 3, 3)$			RSMAL(38)	
E	SU	0	0	0	F		$\Delta H(53), (2, 3, 4)$			RSMAL(53)	
E	0	0	0	0	F		$\Delta H(68), (2, 3, 5)$			RSMAL(68)	
E	SU	0	0	0	0		$\Delta H(11), (2, 4, 1)$			RSMAL(11)	
E	SU	0	0	0	0		$\Delta H(26), (2, 4, 2)$			RSMAL(26)	
E	SU	0	0	0	0		$\Delta H(41), (2, 4, 3)$			RSMAL(41)	
E	SU	0	0	0	0		$\Delta H(56), (2, 4, 4)$			RSMAL(56)	
E	0	0	0	0	0		$\Delta H(71), (2, 4, 5)$			RSMAL(71)	

(B) SSOR SOLUTION MATRICES -- WITH WELL AT NODE SET I=2, J=2 AND OPEN TO ALL 5 LAYERS.

Figure A-3. Example of slice-successive-over-relaxation solution matrices without (A) and with (B) multiaquifer-well coefficients for cross-sectional model consisting of 3 rows, 5 columns, and 5 layers.

putation stencil is positioned at layers 3, 2, and 1, respectively. For the cross-sectional example of figure A-1, the compressed coefficient matrix (array 'A') would have the structure shown in figure A-3B if a well is located in node set (I=2,J=2) and open to all layers.

The computation and loading of multiaquifer-well coefficients for array 'A' are performed in subroutine MULT (appendix G), which is called from subroutine SOLVE prior to development of the standard 'A' array. In subroutine MULT, the coefficients are developed for the entire model before returning to subroutine SOLVE. Because of the symmetry discussed above and illustrated in figures A-2B and C, only those coefficients above any particular location of the computation stencil needs to be explicitly computed. As the E and SU coefficients are computed for the conventional stencil in subroutine SOLVE, the corresponding nonzero elements of the 'A' array generated by subroutine MULT are added to the E and SU terms to form the E' and SU' coefficients.

The multiaquifer-well terms required for the RHS vector (eqs. A-2 and A-3) consisting of coefficients of layers open to a well times the corresponding head prevailing at the end of the previous time step is computed in subroutine RESI (appendix G). This subroutine is called from subroutine SOLVE each time the conventional RHS term is computed. The multiaquifer-well contribution termed (RESO) is added to the RHS after each return to subroutine SOLVE.

Sections of subroutine SOLVE (appendix F) that contain modifications for the multiaquifer-well effects and that call subroutines MULT and RESI are indicated by asterisks in columns 73-75.

Subroutine ALOCAT (appendix G) computes all individual flow components of each node of a multiaquifer node set, including layer flow to the wells. This subroutine is not required for the simulation of multiaquifer-well effects and was written primarily to aid in debugging the multiaquifer-well modification of the SSOR procedure. If the IALOCA option is invoked, subroutine ALOCAT is called from subroutine CHECKI (statement CHK1178).



# APPENDIX B.--PROGRAM VARDEN

```

C *****VAR0010
C *          PROGRAM VARDEN          *VAR0020
C *          -----                  *VAR0030
C * PROGRAM TO COMPUTE CONDUCTANCES AND SOURCE TERMS FOR SIMULATION OF *VAR0040
C * STEADY STATE VARIABLE DENSITY EFFECTS WITH TRECOTT 3-D FLOW MODEL.*VAR0050
C * PROGRAM IS BASED ON R.J.SUN QUASI-3D DEVELOPMENT (WRITTEN COMMUN., *VAR0060
C * 1981) OF G.D.BENNETT VARIABLE-DENSITY FORMULATION (WRITTEN COMMUN. *VAR0070
C * ,1979).                          *VAR0080
C * THE GENERAL STRUCTURE OF THE PROGRAM IS PATTERNED IN PART AFTER *VAR0090
C * PROGRAM WRITTEN BY E.WEISS,1982,(' A MODEL FOR THE SIMULATION OF *VAR0100
C * FLOW OF VARIABLE-DENSITY GROUND WATER IN THREE DIMENSIONS UNDER *VAR0110
C * STEADY STATE CONDITIONS',USGS OPEN FILE REPORT 82-352).          *VAR0120
C * THE PROGRAM MAY BE EXECUTED TO GENERATE DATA FOR TRECOTT MODEL IN *VAR0130
C * TERMS OF FRESH WATER HEADS OR IN TERMS OF PRESSURE. OUTPUT CONSISTS*VAR0140
C * OF ARRAYS WHICH SERVE AS INPUT TO TRECOTT MODEL AND ARRAYS REQUIR-*VAR0150
C * ED TO COMPUTE FLOW DIRECTION AND MAGNITUDE.                      *VAR0160
C *          INPUT                      *VAR0170
C *          -----                  *VAR0180
C * INPUT CONSISTS OF A SINGLE PROGRAM CONTROL RECORD (ON UNIT 5) CON- *VAR0190
C * TAINING THE NUMBER OF MODEL ROWS(IO),COLUMNS(JO), LAYERS(KO); THE *VAR0200
C * LOGICAL UNIT NUMBER OF INPUT ARRAYS(LIN); AND 2 OPTION CODES ----- *VAR0210
C * (IPRESS AND IDELTA). FOLLOWING THIS RECORD THE PROGRAM READS A SER-*VAR0220
C * IES OF INPUT ARRAYS AND THE MODEL GRID SPACING.                  *VAR0230
C * FOR FRESH WATER HEAD SIMULATION IPRESS=0 AND FOR PRESSURE SIMULA- *VAR0240
C * TION IPRESS=1.                                                    *VAR0250
C * THE PROGRAM HAS CODE TO MODIFY TRANSMISSIVITY OF EACH LAYER ,AND *VAR0260
C * THE VERTICAL LEAKANCE BETWEEN LAYERS BY ARRAYS OF MULTIPLICATIVE *VAR0270
C * FACTORS (DELTA ARRAYS). INPUT TRANSMISSIVITY AND VERTICAL LEAKANCE *VAR0280
C * ARE MULTIPLIED BY THE DELTA ARRAYS IF THE OPTION IS INVOKED.      *VAR0290
C * IF MODIFICATION OF VERTICAL LEAKANCE IS DESIRED IDELTA=1 AND IF *VAR0300
C * NOT, IDELTA=0.                                                    *VAR0310
C * TO BYPASS MODIFICATION OF TRANSMISSIVITY AND VERTICAL LEAKANCE SET *VAR0320
C * DELTA ARRAYS TO UNITY AND SET IDELTA=0.                          *VAR0330
C *          *VAR0340
C * INPUT ARRAYS READ FROM UNIT LIN CONSIST OF STORAGE COEFFICIENTS *VAR0350
C * , TRANSMISSIVITY, NODE ELEVATIONS, GROUND WATER DENSITY, *VAR0360
C * PURE WATER DENSITY, VERTICAL LEAKANCE, CONFINING LAYER THICKNESS, *VAR0370
C * VISCOSITY, AND DELTA FACTORS. ALL ARRAYS ARE READ BY SUBROUTINE *VAR0380
C * READS SIMILAR TO THE ARRAY INPUT PROCEEDURE OF THE TRECOTT MODEL. *VAR0390
C * (SEE COMMENTS IN SUBROUTINE READS).                               *VAR0400
C * IT IS TO BE NOTED THAT THE (X,Y) ORIGIN OF THE VARDEN PROGRAM IS *VAR0410
C * AT THE LOWER LEFT CORNER OF EACH LAYER. CONSEQUENTLY INPUT ARRAYS *VAR0420
C * ARE PRESUMED TO RUN FROM BOTTOM ROW TO TOP. LIKEWISE OUTPUT ARRAYS *VAR0430
C * ARE WRITTEN FROM BOTTOM ROW TO TOP. IN THE PROGRAM THE (I+1) INDEX *VAR0440
C * CORRESPONDS TO THE (I-1) INDEX OF THE VARIABLE DENSITY FORMULATION.*VAR0450
C * TO CONFORM TO THE TRECOTT MODEL ROW CONVENTION, MODEL INPUT ARRAYS*VAR0460
C * ARE READ IN REVERSE ROW ORDER IN THE TRECOTT MODEL.              *VAR0470
C * FORMATS OF INPUT DATA AND PRINT OF INPUT DATA ARE CONTROLLED BY *VAR0480
C * INFT AND IOFT ARRAYS PASSED INTO SUBROUTINE READS. SEE DESCRIPTION *VAR0490
C * OF INFT AND IOFT ARRAYS IN DATA STATEMENT.                      *VAR0500
C * IF THE ARGUMENT OF INFT IS (1,1) THE INPUT FORMAT IS 10F8.0; IF THE*VAR0501

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# APPENDIX B (Continued)

```

C * ARGUMENT IS (1,2) THE FORMAT IS 10E8.2; AND IF THE ARGUMENT IS      *VAR0502
C * (1,3) THE FORMAT IS 20F4.0. FOR PRINTED OUTPUT THE ARGUMENT OF(1,4)*VAR0503
C * OF ARRAY IOFT PASSED TO SUBROUTINE READS WILL RESULT IN PRINTED      *VAR0504
C * ARRAYS IN 10E11.3 FORMAT.                                           *VAR0505
C *                                                                      *VAR0510
C * IT IS ASSUMED THAT INPUT DENSITIES ARE IN KILOGRAMS/CUBIC METER.    *VAR0520
C * WITHIN PROGRAM, WHEN PURE WATER DENSITY IS READ IT IS CONVERTED TO *VAR0530
C * GRAMS/CUBIC CENTIMETER, AND THEN CONVERTED TO A RELATIVE DENSITY BY*VAR0540
C * DIVIDING BY THE DENSITY OF FRESHWATER (TAKEN TO BE 0.9977).         *VAR0550
C * THUS, PURE WATER DENSITY IS REPLACED BY 0.9977*RELATIVE DENSITY.    *VAR0560
C * IF TERMS FOR PRESSURE ARE TO BE DEVELOPED, CONDUCTANCES ARE DIVIDED*VAR0570
C * BY ACCELERATION OF GRAVITY (32.174) RATHER THAN 1.94*.9977*32.174. *VAR0580
C * ALSO RELATIVE GROUND WATER DENSITY IS MULTIPLIED BY 1.94*.9977*G    *VAR0590
C * FOR PRESSURE FORMULATION.                                           *VAR0600
C * TO CONVERT DENSITIES AND VISCOSITY TO RELATIVE DENSITY AND RELATIVE*VAR0610
C * VISCOSITY USE FAC MULTIPLIER OF ARRAY CONTROL RECORD.              *VAR0620
C * FOR NORTHERN MIDWEST RASA MODEL USED FAC=1/(.9977) FOR DENSITY AND *VAR0630
C * FAC=1/(1.127 CENTIPOISE)) FOR VISCOSITY.                           *VAR0640
C * NOTE THAT THE DENSITY OF 0.9977 and VISCOSITY OF 1.127 REPRESENT   *VAR0650
C * AMBIENT VALUES FOR FRESHWATER AREAS OF THE NORTHERN MIDWEST       *VAR0655
C * FOLLOWING INPUT OF THE ABOVE ARRAYS, THE MODEL GRID SPACING ALONG  *VAR0660
C * ROWS (DELX), AND COLUMNS (DELY) IS READ FROM UNIT 5 IN SUBROUTINE *VAR0670
C * DEL (SEE COMMENTS IN SUBROUTINE DEL).                               *VAR0680
C *                                                                      *VAR0690
C *                               OUTPUT                                  *VAR0700
C * -----
C * OUTPUT OF VARDEN PROGRAM ARRAYS IS IMPLEMENTED IN SUBROUTINE OUTP.  *VAR0710
C * OUTPUT ARRAYS AND THEIR PROGRAM UNIT NUMBERS ARE AS FOLLOWS:      *VAR0720
C * MODIFIED TRANSMISSIVITY (TOUT)-- UNIT 18                           *VAR0730
C * MODIFIED STORAGE COEFFICIENTS (SOUT)-- UNIT 19                     *VAR0740
C * CONDUCTANCE ALONG ROWS (TR)--UNIT 11                               *VAR0750
C * CONDUCTANCE ALONG COLUMNS (TC)--UNIT 12                           *VAR0760
C * INDIVIDUAL SOURCE TERM FOR FLOW ALONG ROWS (ALPHA)--UNIT 13        *VAR0770
C * INDIVIDUAL SOURCE TERM FOR FLOW ALONG COLUMNS (BETA)--UNIT 14     *VAR0780
C * TOTAL SOURCE TERM (PST)--UNIT 16                                   *VAR0790
C * INDIVIDUAL SOURCE TERM FOR FLOW BETWEEN LAYERS (GAMMA)--UNIT 17    *VAR0800
C * CONDUCTANCE BETWEEN LAYERS (TK)--UNIT 15                           *VAR0810
C * (SEE SUBROUTINE OUTP FOR FORMAT OF OUTPUT ARRAYS).                 *VAR0820
C * PRINT OUTPUT OF PROGRAM IS WRITTEN TO UNIT 6.                       *VAR0830
C *                                                                      *VAR0840
C * IN ADDITION TO THE ABOVE INPUT AND OUTPUT ARRAYS, INTERNAL PROGRAM *VAR0850
C * ARRAYS INCLUDE PSA,PSB,AND PS3 TERMS OF EQ. 5C.                   *VAR0860
C *                                                                      *VAR0865
C * DIMENSIONING OF PROGRAM IS SET BY PARAMETER STATEMENT, WHERE IR IS *VAR0870
C * NUMBER OF MODEL ROWS, IC IS NUMBER OF MODEL COLUMNS, AND KL IS NUM-*VAR0880
C * BER OF MODEL LAYERS.                                               *VAR0890
C * SUBROUTINES REQUIRED BY VARDEN ARE:                                *VAR0900
C * SUBROUTINE READS TO READ INPUT ARRAYS (OTHER THAN GRID SPACING)    *VAR0910
C * SUBROUTINE DEL TO READ GRID SPACING                                *VAR0920
C * SUBROUTINE ZERO TO SET ARRAYS TO ZERO AFTER COMPUTATIONS FOR LAYER *VAR0930
C * K ARE COMPLETED                                                  *VAR0940
C * SUBROUTINE OUTP TO WRITE OUTPUT ARRAYS IN TO SPECIFIED FILES.      *VAR0950
C * DIMENSIONING OF SUBROUTINES IS SET BY PROGRAM.                    *VAR0960

```

# APPENDIX B (Continued)

```

C * *****VAR0970
C
PARAMETER IR=37,IC=40,KL=5,KV=KL+1,KE=KL+2
DIMENSION
+ ALPHA(IR,IC), BETA(IR,IC), DELTA(IR,IC,KL),
+ DELX(IC), DELY(IR), DEN(IR,IC,KE),
+ DENP(IR,IC,KE), ELEV(IR,IC,KL), GAMMA(IR,IC),
+ INFT(2,3), IOFT(9,4), PS3(IR,IC),
+ PSA(IR,IC), PSB(IR,IC), PST(IR,IC),
+ S(IR,IC,KL), T(IR,IC,KL), TC(IR,IC),
+ THK(IR,IC,KV), TK(IR,IC), TKV(IR,IC,KV),
+ TOUT(IR,IC), TR(IR,IC), VISC(IR,IC,KE)
+ ,SOUT(IR,IC)
COMMON /MISC/
+ IO, JO
DATA INFT/4H(10F,4H8.0),4H(10E,4H8.2),4H(20F,4H4.0)/
DATA IOFT/4H(1H0,4H,I2,,4H2X,2,4HOF6.,4H1/(5,4HX,20,4HF6.1,4H)) ,
+4H ,4H(1H0,4H,I5,,4H14F9,4H.5/(,4H1H ,4H5X,1,4H4F9.,4H5)) ,4H
+ ,4H(1H0,4H,I5,,4H10E1,4H2.5/,4H(1H ,4H,5X,,4H10E1,4H2.5),4H)
+ ,4H(1H0,4H,I5,,4H10E1,4H1.3/,4H(1H ,4H,5X,,4H10E1,4H1.3),4H) /
C
C READ IN VARDEN PROGRAM CONTROL RECORD.
C
READ (5,1) IO,JO,KO,LIN,IPRESS,IDELTA
1 FORMAT(6I5)
WRITE (6,2) IO,JO,KO,IPRESS
2 FORMAT (1H0,4I5)
IF(IDELTA.NE.0) WRITE(6,986)
986 FORMAT(1H0,' ***** DELTA FACTOR APPLIED TO TK *****')
IF (IPRESS.EQ.0) WRITE (6,3)
3 FORMAT (1H0,'***** COEFFICIENTS FOR FRESHWATER HEAV
+D*****')
C
DDELTU=1.0
DDELTL=1.0
KVV=KO+1
KEE=KO+2
K1=KO-1
J2=JO-2
I2=IO-2
C
C NOTE IN CALL TO READS PASS APPROPRIATE INDEX OF LAYERS.
C READ STORAGE COEFFICIENTS
C
DO 4 K=1,KO
4 CALL READS (S(1,1,K),INFT(1,2),IOFT(1,4),'S ARRAYS
+,LIN,K)
C
C READ TRANSMISSIVITY
C

```

# APPENDIX B (Continued)

	DO 5 K=1,KO	VAR1470
	5 CALL READS (T(1,1,K),INFT(1,2),IOFT(1,4),'T ARRAYS	'VAR1480
	+ ,LIN,K)	VAR1490
C		VAR1500
C	READ IN NODE ELEVATIONS	VAR1510
C		VAR1520
	DO 6 K=1,KO	VAR1530
	CALL READS (ELEV(1,1,K),INFT(1,1),IOFT(1,4),'ELEVATION ARRAYS	VAR1540
	+ ' ,LIN,K)	VAR1550
	6 CONTINUE	VAR1560
C		VAR1570
C	READ DENSITY OF SOLUTION	VAR1580
C		VAR1590
	DO 7 K=2,KVV	VAR1600
	7 CALL READS (DEN(1,1,K),INFT(1,1),IOFT(1,4),'RHO-S ARRAYS	'VAR1610
	+ ,LIN,K)	VAR1620
C		VAR1630
C	READ DENSITY OF PURE WATER	VAR1640
C		VAR1650
	DO 8 K=2,KVV	VAR1660
	8 CALL READS (DENP(1,1,K),INFT(1,1),IOFT(1,4),'DEN-P ARRAYS	VAR1670
	+ ' ,LIN,K)	VAR1680
C		VAR1690
C	READ VERTICAL LEAKANCE	VAR1700
C		VAR1710
	DO 9 K=2,KO	VAR1720
	9 CALL READS (TKV(1,1,K),INFT(1,2),IOFT(1,4),'LEAKANCE ARRAYS	VAR1730
	+ ' ,LIN,K)	VAR1740
C		VAR1750
C	READ THICKNESS OF CONFINING LAYERS	VAR1760
C		VAR1770
	DO 10 K=2,KO	VAR1780
	10 CALL READS (THK(1,1,K),INFT(1,1),IOFT(1,4),'CONFINING BED THICKNES	VAR1790
	+S' ,LIN,K)	VAR1800
C		VAR1810
C	READ IN VISCOSITY	VAR1820
C		VAR1830
	DO 11 K=2,KVV	VAR1840
	11 CALL READS (VISC(1,1,K),INFT(1,1),IOFT(1,4),'VISCOSITY ARRAY	VAR1850
	+ ' ,LIN,K)	VAR1860
C		VAR1870
C	READ IN PERMEABILITY/OVERBURDEN FACTOR	VAR1880
C		VAR1890
	DO 12 K=1,KO	VAR1900
	12 CALL READS (DELTA(1,1,K),INFT(1,1),IOFT(1,4),'K VARIATION WITH DEP	VAR1910
	+TH ' ,LIN,K)	VAR1920
	I1=I0-1	VAR1930
	J1=J0-1	VAR1940
C		VAR1950
C	FOR CROSS-SECTIONAL MODEL IN J DIRECTION(X DIRECTION) SET	VAR1960
C	BOUNDARY ELEVATIONS ALONG I=1,AND I=3 TO BE THE SAME AS	VAR1970
C	ELEVATIONS OF NODES ALONG CROSS-SECTION.	VAR1980

# APPENDIX B (Continued)

C		VAR1990
	IF (I0.GT.3) GO TO 16	VAR2000
	DO 15 K=1,K0	VAR2010
	DO 14 J=1,J0	VAR2020
	ELEV(1,J,K)=ELEV(2,J,K)	VAR2030
	ELEV(3,J,K)=ELEV(2,J,K)	VAR2040
14	CONTINUE	VAR2050
15	CONTINUE	VAR2060
16	CONTINUE	VAR2070
C		VAR2080
C	*****	VAR2090
C	THE FOLLOWING COMMENTED STATEMENTS PERTAIN TO CODE APPLICABLE TO	VAR2100
C	NORTHERN-MIDWEST MODEL. FOR NODES IN "PINCHOUT" AREAS T,DELTA,VISC,	VAR2110
C	DENP,AND DEN ARE SET TO CORRESPONDING VALUES OF TOP LAYER. ALSO SET	VAR2120
C	TKV OF CONFINING BEDS BOUNDED BY INACTIVE NODES TO ZERO.	VAR2130
C		VAR2140
C	DO 19 K=1,K1	VAR2150
C	DO 18 I=1,I0	VAR2160
C	DO 17 J=1,J0	VAR2170
C	TT=T(I,J,K)	VAR2180
C	SS=S(I,J,K)	VAR2190
C	IF(SS.LT.0.0.AND.S(I,J,K+1).LT.0.0.AND.TT.EQ.0.0) TKV(I,J,K+1)=0.0	VAR2200
C	IF ((TT.EQ.0.0).AND.(SS.LT.0.0)) DELTA(I,J,K)=DELTA(I,J,K0)	VAR2210
C	IF ((TT.EQ.0.0).AND.(SS.LT.0.0)) VISC(I,J,K+1)=VISC(I,J,KV)	VAR2220
C	IF ((TT.EQ.0.0).AND.(SS.LT.0.0)) DENP(I,J,K+1)=DENP(I,J,KV)	VAR2230
C	IF ((TT.EQ.0.0.AND.(SS.LT.0.0)) DEN(I,J,K+1)=DENP(I,J,KV)	VAR2240
C	17 IF ((T(I,J,K).EQ.0.0).AND.(S(I,J,K).LT.0.0)) T(I,J,K)=T(I,J,K0	VAR2250
C	+ )	VAR2260
C	18 CONTINUE	VAR2270
C	19 CONTINUE	VAR2280
C	*****	VAR2290
C		VAR2300
C	ZERO BORDER NODES	VAR2310
C		VAR2320
	DO 22 K=1,K0	VAR2330
	DO 20 I=1,I0	VAR2340
	S(I,1,K)=-1.0	VAR2350
	T(I,1,K)=0.0	VAR2360
	S(I,J0,K)=-1.0	VAR2370
	T(I,J0,K)=0.0	VAR2380
20	CONTINUE	VAR2390
	DO 21 J=2,J1	VAR2400
	S(1,J,K)=-1.0	VAR2410
	T(1,J,K)=0.0	VAR2420
	S(I0,J,K)=-1.0	VAR2430
	T(I0,J,K)=0.0	VAR2440
21	CONTINUE	VAR2450
22	CONTINUE	VAR2460
C		VAR2470
C	ZERO DUMMY ARRAYS AT TOP AND BOTTOM OF MODEL.	VAR2480

# APPENDIX B (Continued)

C		VAR2490
	KA=KVV	VAR2500
	KB=KEE	VAR2510
	DO 24 I=1,I0	VAR2520
	DO 23 J=1,J0	VAR2530
	TKV(I,J,1)=0.0	VAR2540
	TKV(I,J,KA)=0.0	VAR2550
	THK(I,J,1)=0.0	VAR2560
	THK(I,J,KA)=0.0	VAR2570
	DEN(I,J,1)=0.0	VAR2580
	DEN(I,J,KB)=0.0	VAR2590
	DENP(I,J,1)=0.0	VAR2600
	DENP(I,J,KB)=0.0	VAR2610
	VISC(I,J,1)=0.0	VAR2620
	VISC(I,J,KB)=0.0	VAR2630
23	CONTINUE	VAR2640
24	CONTINUE	VAR2650
C		VAR2660
C	READ IN DELX,DELY	VAR2670
C		VAR2680
	CALL DEL (DELX,1,J0)	VAR2690
	CALL DEL (DELY,2,I0)	VAR2700
C		VAR2710
C	SET UP UNIT FACTORS FOR FRESHWATER OR PRESSURE FORMULATION	VAR2720
C	IN BRITISH SYSTEM OF UNITS	VAR2730
C		VAR2740
	UNIT1=1.94*.9977	VAR2750
	UNIT2=1.0	VAR2760
	IF (IPRESS.NE.0) UNIT1=1.0/32.174	VAR2770
	IF (IPRESS.NE.0) UNIT2=32.174*1.94*.9977	VAR2780
	SUMA=0.0	VAR2790
	SUMB=0.0	VAR2800
	SUMC=0.0	VAR2810
	SUMT=0.0	VAR2820
C		VAR2830
C	MODIFY TRANSMISSIVITY TO INCORPORATE DELTA, DENSITY	*VAR2840
C	AND VISCOSITY.	*VAR2850
C		VAR2860
	DO 27 K=1,K0	VAR2870
	DO 26 I=2,I1	VAR2880
	DO 25 J=2,J1	VAR2890
25	T(I,J,K)=T(I,J,K)*DELTA(I,J,K)*DENP(I,J,K+1)/VISC(I,J,K+1)	VAR2900
26	CONTINUE	VAR2910
27	CONTINUE	VAR2920
C		VAR2930
C	***** START OF MAJOR LOOP *****	VAR2940
C	ZERO ARRAYS AFTER COMPUTATION FOR LAYER K	VAR2950
C		VAR2960
	DO 42 K=1,K0	VAR2970
	CALL ZERO (TR)	VAR2980
	CALL ZERO (TC)	VAR2990
	CALL ZERO (PSA)	VAR3000

# APPENDIX B (Continued)

	CALL ZERO (PSB)	VAR3010
	CALL ZERO (PS3)	VAR3020
	CALL ZERO (PST)	VAR3030
	CALL ZERO (ALPHA)	VAR3040
	CALL ZERO (BETA)	VAR3050
	CALL ZERO (GAMMA)	VAR3060
	CALL ZERO (TOUT)	VAR3070
	CALL ZERO(SOUT)	VAR3080
	CALL ZERO(TK)	VAR3090
	SMA=0.0	VAR3100
	SMB=0.0	VAR3110
	SMG=0.0	VAR3120
	SMT=0.0	VAR3130
	DO 30 I=I1,2,-1	VAR3140
	DO 29 J=2,J1	VAR3150
	IF (T(I,J,K).EQ.0.0) GO TO 29	VAR3160
C		VAR3170
C	COMPUTE GEOMETRY OF MODEL	VAR3180
C		VAR3190
	E=ELEV(I,J,K)	VAR3200
	DZAP=ELEV(I,J+1,K)-E	VAR3210
	DZAM=-ELEV(I,J-1,K)+E	VAR3220
	DZBP=-ELEV(I+1,J,K)+E	VAR3230
	DZBM=ELEV(I-1,J,K)-E	VAR3240
	DJP=DELX(J+1)/2.	VAR3250
	DJ=DELX(J)/2.	VAR3260
	DJM=DELX(J-1)/2.	VAR3270
	DIP=DELY(I+1)/2.	VAR3280
	DI=DELY(I)/2.	VAR3290
	DIM=DELY(I-1)/2.	VAR3300
	DAP=SQRT((DJP+DJ)**2.+DZAP**2.)	VAR3310
	DAM=SQRT((DJM+DJ)**2.+DZAM**2.)	VAR3320
	DBP=SQRT((DIP+DI)**2.+DZBP**2.)	VAR3330
	DBM=SQRT((DIM+DI)**2.+DZBM**2.)	VAR3340
	DA=2.0*DJ*DAM/(DJM+DJ)	VAR3350
	DA2=2.0*DJP*DAP/(DJ+DJP)	VAR3360
	DB=2.0*DI*DBM/(DIM+DI)	VAR3370
	DB1=2.0*(DBM-DB/2.)	VAR3380
	DB2=2.0*DIP*DBP/(DI+DIP)	VAR3390
C		VAR3400
C	DEVELOP HARMONIC MEANS FOR PRESSURE OR HEAD FORMULATION	VAR3410
C		VAR3420
C	COMPUTE AND STORE HORIZONTAL CONDUCTANCE FOR OUTPUT	VAR3430
	T1=T(I,J,K)	VAR3440
	T2=T(I,J+1,K)	VAR3450
	T3=T(I-1,J,K)	VAR3460
	TR(I,J)=((2.*T1*DB)*(T2*UNIT1))/(T2*DA+T1*DA2)	VAR3470
	TC(I,J)=((2.*T1*DA)*(T3*UNIT1))/(T3*DB+T1*DB1)	VAR3480
C		VAR3490
C		VAR3500

# APPENDIX B (Continued)

C		VAR3510
C	COMPUTE AND STORE STORAGE TERMS FOR SUBSEQUENT OUTPUT	VAR3520
C		VAR3530
	IF(S(I,J,K).EQ.-1.) SOUT(I,J)=-1.0	VAR3540
	IF(S(I,J,K).EQ.-1.0) GO TO 222	VAR3550
	SOUT(I,J)=S(I,J,K)*DA*DB*UNIT1	VAR3560
	222 CONTINUE	VAR3570
C		VAR3580
C	DEVELOPE VERTICAL CONDUCTANCE COEFFICIENTS	VAR3590
C	NOTE THAT M+1 INDEX CORRESPONDS TO LAYER K INDEX	VAR3600
C		VAR3610
	M=K	VAR3620
	DDENPL=(DENP(I,J,M+1)+DENP(I,J,M))/2.0	VAR3630
	DDENPU=(DENP(I,J,M+1)+DENP(I,J,M+2))/2.0	VAR3640
	DVISCL=(VISC(I,J,M+1)+VISC(I,J,M))/2.0	VAR3650
	DVISCU=(VISC(I,J,M+1)+VISC(I,J,M+2))/2.0	VAR3660
C		VAR3670
C	MODIFY VERTICAL CONDUCTANCE BY DELTA FACTOR IF IDELTA IS	VAR3680
C	NOT ZERO.	*VAR3690
C		VAR3700
	IF (IDELTA.EQ.0) GO TO 28	VAR3710
	IF (K.NE.K0) DDELTU=(DELTA(I,J,K)+DELTA(I,J,K+1))*0.5	VAR3720
	IF (K.NE.1) DDELTU=(DELTA(I,J,K)+DELTA(I,J,K-1))*0.5	VAR3730
	IF (K.EQ.1) DDELTU=0.0	VAR3740
	IF (K.EQ.K0) DDELTU=0.0	VAR3750
28	CONTINUE	VAR3760
	TZL=TKV(I,J,M)*DA*DB*DDENPL*UNIT1*DDELTU/DVISCL	VAR3770
	TZU=TKV(I,J,M+1)*DA*DB*DDENPU*UNIT1*DDELTU/DVISCU	VAR3780
C		VAR3790
C	STORE COEFFICIENTS FOR MODEL INPUT	VAR3800
C		VAR3810
	IF (K.LT.K0) TK(I,J)=TZU	VAR3820
C		VAR3830
C	DEVELOP HORIZONTAL SOURCE TERMS	VAR3840
C		VAR3850
	M=K+1	VAR3860
	D=DEN(I,J,M)	VAR3870
	DNAP=((DEN(I,J+1,M)+D)/2.0)*UNIT2	VAR3880
	DNAM=((DEN(I,J-1,M)+D)/2.0)*UNIT2	VAR3890
	DNBP=((DEN(I+1,J,M)+D)/2.0)*UNIT2	VAR3900
	DNBM=((DEN(I-1,J,M)+D)/2.0)*UNIT2	VAR3910
C		VAR3920
C	ALTER DENSITY TERM IF FRESH WATER FORMULATION	VAR3930
C		VAR3940
	IF (IPRESS.EQ.0) DNAP=(DNAP-1.0)	VAR3950
	IF (IPRESS.EQ.0) DNAM=(DNAM-1.0)	VAR3960
	IF (IPRESS.EQ.0) DNBP=(DNBP-1.0)	VAR3970
	IF (IPRESS.EQ.0) DNBM=(DNBM-1.0)	VAR3980
	ALPHA(I,J)=-TR(I,J)*DNAP*DZAP	VAR3990
	BETA(I,J)=-TC(I,J)*DNBM*DZBM	VAR4000
	PSA(I,J)=-TR(I,J)*DNAP*DZAP+TR(I,J-1)*DNAM*DZAM	VAR4010
	PSB(I,J)=-TC(I,J)*DNBM*DZBM+TC(I+1,J)*DNBP*DZBP	VAR4020



# APPENDIX B (Continued)

	SUMA=PSA(I,J)+SUMA	VAR4030
	SUMB=PSB(I,J)+SUMB	VAR4040
C		VAR4050
C	DEVELOP VERTICAL SOURCE TERM	VAR4060
C		VAR4070
	M=K	VAR4080
	DDENL=(DEN(I,J,M+1)+DEN(I,J,M))/2.0*UNIT2	VAR4090
	DDENU=(DEN(I,J,M+1)+DEN(I,J,M+2))/2.0*UNIT2	VAR4100
C		VAR4110
C	ALTER DENSITY TERM IF FRESHWATER FORMULATION	VAR4120
C		VAR4130
	IF (IPRESS.EQ.0) DDENL=DDENL-1.0	VAR4140
	IF (IPRESS.EQ.0) DDENU=DDENU-1.0	VAR4150
	IF (THK(I,J,K+1).LT.1.0) THK(I,J,K+1)=1.0	VAR4160
	IF (THK(I,J,K).LT.1.0) THK(I,J,K)=1.0	VAR4170
	GAMMA(I,J)=-TZU*DDENU*THK(I,J,K+1)*(DELX(J)*DELY(I)/(DA*DB))	VAR4180
	PS3(I,J)=(+TZL*DDENL*THK(I,J,K)-TZU*DDENU*THK(I,J,K+1))*(DEL	VAR4190
+	X(J)*DELY(I)/(DA*DB))	VAR4200
	PST(I,J)=-PSA(I,J)-PSB(I,J)-PS3(I,J)	VAR4210
	SUMC=PS3(I,J)+SUMC	VAR4220
	SUMT=PST(I,J)+SUMT	VAR4230
	SMA=PSA(I,J)+SMA	VAR4240
	SMB=PSB(I,J)+SMB	VAR4250
	SMG=PS3(I,J)+SMG	VAR4260
	SMT=PST(I,J)+SMT	VAR4270
29	CONTINUE	VAR4280
30	CONTINUE	VAR4290
C		VAR4300
C	*****	VAR4310
C	THE FOLLOWING COMMENTED STATEMENTS PERTAIN TO CODE APPLICABLE TO	VAR4320
C	NORTHERN-MIDWEST MODEL. FOR INACTIVE NODES CONDUCTANCES AND SOURCE	VAR4330
C	TERMS ARE SET TO ZERO.	VAR4340
C		VAR4350
C	IF (K.EQ.K0) GO TO 35	VAR4360
C	DO 32 I=I1,2,-1	VAR4370
C	DO 31 J=2,J2	VAR4380
C	T5=T(I,J,K0)	VAR4390
C	IF (T(I,J,K).NE.T5) GO TO 31	VAR4400
C	IF ((T(I,J+1,K).NE.T5).AND.(S(I,J+1,K).GE.0.0)) GO TO 31	VAR4410
C	ALPHA(I,J)=0.0	VAR4420
C	PST(I,J)=PST(I,J)+PSA(I,J)	VAR4430
C	TR(I,J)=0.0	VAR4440
C	31 CONTINUE	VAR4450
C	32 CONTINUE	VAR4460
C	DO 34 J=2,J1	VAR4470
C	DO 33 I=I1,3,-1	VAR4480
C	T5=T(I,J,K0)	VAR4490
C	IF (T(I,J,K).NE.T5) GO TO 33	VAR4500
C	IF ((T(I-1,J,K).NE.T5).AND.(S(I-1,J,K).GE.0.0)) GO TO 33	VAR4510
C	BETA(I,J)=0.0	VAR4520
C	PST(I,J)=PST(I,J)+PSB(I,J)	VAR4530
C	TC(I,J)=0.0	VAR4540

# APPENDIX B (Continued)

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C 33      CONTINUE                                VAR4550
C 34      CONTINUE                                VAR4560
C 35      CONTINUE                                VAR4570
C *****VAR4580
C                                VAR4590
C          SET TRANSMISSIVITY TO CONFORM TO MULTIAQUIFER FORMULATION. VAR4600
C                                VAR4610
C          DO 37 I=2,I1                            VAR4620
C            DO 36 J=2,J1                            VAR4630
36      TOUT(I,J)=T(I,J,K)*UNIT1                    VAR4640
37      CONTINUE                                    VAR4650
C  PRINT SUMMATION OF SOURCE TERMS FOR LAYER K      VAR4660
C          WRITE (6,38)                            VAR4670
38      FORMAT (1H0,125('*'))                      VAR4680
C          WRITE (6,39) K                          VAR4690
39      FORMAT (1X,'LAYER',1X,I1,5X,'SUMA,SUMB,SUMG,SUMT') VAR4700
C          WRITE (6,40) SMA,SMB,SMG,SMT            VAR4710
40      FORMAT (1X,4(E12.5))                        VAR4720
C          OUTPUT ARRAYS FOR LAYER K                VAR4730
C          CALL OUTP (TOUT,K,18,'MULTIAQUIFER T      ',2) VAR4740
C          CALL OUTP (SOUT,K,19,'STORAGE COEFFICIENTS ',1) VAR4750
C          CALL OUTP (TR,K,11,'TR COEFFICIENTS        ',2) VAR4760
C          CALL OUTP (TC,K,12,'TC COEFFICIENTS        ',2) VAR4770
C          CALL OUTP (ALPHA,K,13,'ALPHA SOURCE TERM    ',2) VAR4780
C          CALL OUTP (BETA,K,14,'BETA SOURCE TERM      ',2) VAR4790
C          CALL OUTP (PST,K,16,'TOTAL SOURCE TERM      ',2) VAR4800
C          IF (K.EQ.K0) GO TO 41                    VAR4810
C          CALL OUTP (GAMMA,K,15,'GAMMA SOURCE TERM    ',2) VAR4820
C          CALL OUTP (TK,K,17,'TK COEFFICIENT          ',2) VAR4830
41      CONTINUE                                    VAR4840
42      CONTINUE                                    VAR4850
C  PRINT SUMMATION OF SOURCE TERMS OVER ALL LAYERS  VAR4860
C          WRITE (6,43)                            VAR4870
43      FORMAT (1H0,' SOURCE TERM TOTALS: ALPHA,BETA,GAMMA,TOTAL') VAR4880
C          WRITE (6,44) SUMA,SUMB,SUMC,SUMT          VAR4890
44      FORMAT (1H0,4E12.3)                        VAR4900
C          WRITE (6,45)                            VAR4910
45      FORMAT (1H0,' PROGRAM COMPLETE')            VAR4920
C          STOP                                     VAR4930
C          END                                     VAR4940
C *****VAR4950
C          SUBROUTINE DEL (D,IM,LL)                  VAR4960
C *****VAR4970
C * SUBROUTINE TO READ MODEL NODE SPACING ARRAY IN EITHER ROW OR COLUMN*VAR4980
C * DIRECTION. EACH ARRAY REQUIRES A CONTROL RECORD CONTAINING VALUES *VAR4990
C * OF FAC,IVAR,IPRN. CONTROL RECORD AND NODE SPACING ARRAY ARE READ *VAR5000
C * FROM UNIT 5. SEE COMMENTS IN SUBROUTINE READS CONCERNING FORM OF *VAR5010
C * CONTROL RECORD.                                *VAR5020
C *****VAR5030

```

# APPENDIX B (Continued)

	DIMENSION	VAR5040
	+ D(100)	VAR5050
C		VAR5060
C		VAR5070
	IC=IM	VAR5080
	READ (5,1) FAC,IVAR,IPRN	VAR5090
1	FORMAT (F10.0,2I10)	VAR5100
	IF (IVAR.EQ.1) READ (5,2) (D(J),J=1,LL)	VAR5110
2	FORMAT (8G10.0)	VAR5120
	DO 4 J=1,LL	VAR5130
	IF (IVAR.NE.1) GO TO 3	VAR5140
	D(J)=D(J)*FAC	VAR5150
	GO TO 4	VAR5160
3	D(J)=FAC	VAR5170
4	CONTINUE	VAR5180
	IF (IC.NE.1) GO TO 7	VAR5190
	IF (IVAR.EQ.1.AND.IPRN.NE.1) WRITE (6,5) (D(J),J=1,LL)	VAR5200
5	FORMAT (1H1,'GRID SPACING IN X DIRECTION'//12F10.0)	VAR5210
	IF (IVAR.EQ.0) WRITE (6,6) FAC	VAR5220
6	FORMAT (1H0,72X,'DELX =' ,G15.7)	VAR5230
	RETURN	VAR5240
7	IF (IVAR.EQ.1.AND.IPRN.NE.1) WRITE (6,8) (D(J),J=1,LL)	VAR5250
8	FORMAT (1H1,'GRID SPACING IN Y DIRECTION'//12F10.0)	VAR5260
	IF (IVAR.EQ.0) WRITE (6,9) FAC	VAR5270
9	FORMAT (1H0,72X,'DELY =' ,G15.7)	VAR5280
	RETURN	VAR5290
	END	VAR5300
C*****		VAR5310
	SUBROUTINE OUTP (A,K,LOUT,IN,KOUT)	VAR5320
C*****		VAR5330
C *	SUBROUTINE TO OUTPUT VARDEN PROGRAM ARRAYS IN FORMAT OF TRESCOTT	*VAR5340
C *	MODEL INPUT.	*VAR5350
C *	OUTPUT FORMAT GOVERNED BY VALUE OF KOUT IN SUBROUTINE ARGUMENT	*VAR5360
C *	LIST. IF KOUT =1 FORMAT IS (8F10.0), IF KOUT=2 FORMAT IS (8E10.4),	*VAR5370
C *	IF KOUT=3 FORMAT IS (8F8.4).	*VAR5380
C*****		VAR5390
	CHARACTER*4	VAR5400
	+ IN(6)	VAR5410
	COMMON /MISC/	VAR5420
	+ IO, JO	VAR5430
	DIMENSION	VAR5440
	+ V(2)	VAR5450
	DIMENSION	VAR5460
	+ A(IO,JO)	VAR5470
	DATA (V(I),I=1,2)/4H(10F,4H8.5)/	VAR5480
C		VAR5490
C		VAR5500
	WRITE (6,1) IN,K	VAR5510
1	FORMAT (1H0,2X,6A4,' FOR LAYER',I3/2X,39('-'))	VAR5520
	V(1)=4H(8F1	VAR5530

# APPENDIX B (Continued)

```

IF (KOUT.EQ.2) V(1)=4H(8E1                                VAR5540
V(2)=4H0.4)                                                VAR5550
IF (KOUT.EQ.1) V(2)=4H0.0)                                VAR5560
IF (KOUT.EQ.3) V(2)=4H8.4)                                VAR5570
DO 2 I=1,I0                                                VAR5580
2 WRITE (LOUT,V) (A(I,J),J=1,J0)                          VAR5590
RETURN                                                    VAR5600
END                                                        VAR5610
C*****VAR5620
SUBROUTINE READS (A,INFT,IOFT,IN,LIN,K)                    VAR5630
C *****VAR5640
C * SUBROUTINE READS IS A MODIFICATION OF ENTRY ARRAY OF SUBROUTINE *VAR5650
C * DATAI OF TRESCOTT MODEL. EACH INPUT ARRAY REQUIRES A CONTROL RECORD*VAR5660
C * CONTAINING VALUES OF FAC, IVAR, AND IPRN SIMILAR TO THE CONTROL *VAR5670
C * RECORD FOR ARRAY INPUT OF TRESCOTT MODEL. IF ARRAY IN QUESTION *VAR5680
C * CONSISTS OF A CONSTANT VALUE FOR EACH NODE, SET FAC= THE CONSTANT *VAR5690
C * AND SET IVAR=0. IN THIS CASE NO ARRAY IS READ. *VAR5700
C * IF THE ARRAY VALUES ARE NOT CONSTANT SET FAC TO A CONSTANT AND SET *VAR5710
C * IVAR=1. IN THIS CASE AN ARRAY WILL BE READ AND ALL VALUES ARE MUL- *VAR5720
C * TIPLIED BY FAC. IN EITHER CASE SET IPRN=0 TO PRINT INPUT ARRAY OR *VAR5730
C * SET IPRN=1 IF PRINT OF INPUT ARRAY IS NOT DESIRED. *VAR5740
C *****VAR5750
CHARACTER*4                                                VAR5760
+ IN(6)                                                    VAR5770
COMMON /MISC/                                              VAR5780
+ IO, J0                                                    VAR5790
DIMENSION                                                  VAR5800
+ A(I0,J0), INFT(2,2), IOFT(9,4)                          VAR5810
C                                                         VAR5820
C                                                         VAR5830
READ (5,1) FAC,IVAR,IPRN                                  VAR5840
1 FORMAT (G10.0,2I10,3F10.0,2I10)                        VAR5850
IRECS=0                                                    VAR5860
IC=4*IRECS+2*IVAR+IPRN+1                                  VAR5870
GO TO (2,2,5,5,9,9), IC                                  VAR5880
2 DO 3 I=1,I0                                              VAR5890
DO 3 J=1,J0                                                VAR5900
3 A(I,J)=FAC                                              VAR5910
WRITE (6,4) IN,FAC,K                                       VAR5920
4 FORMAT (1H0,52X,6A4,' =',G15.7,' FOR LAYER',I3)        VAR5930
GO TO 11                                                    VAR5940
5 IF (IC.EQ.3) WRITE (6,6) IN,K                            VAR5950
6 FORMAT (1H0,45X,6A4,' MATRIX, LAYER',I3/46X,41('-'))    VAR5960
DO 8 I=1,I0                                                VAR5970
READ (LIN,INFT) (A(I,J),J=1,J0)                          VAR5980
DO 7 J=1,J0                                                VAR5990
A(I,J)=A(I,J)*FAC                                         VAR6000
7 CONTINUE                                                 VAR6010
8 IF (IC.EQ.3) WRITE (6,IOFT) I,(A(I,J),J=1,J0)          VAR6020
GO TO 11                                                    VAR6030
9 CONTINUE                                                 VAR6040

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# APPENDIX B (Continued)

IF (IC.EQ.6) GO TO 11	VAR6050
WRITE (6,6) IN,K	VAR6060
DO 10 I=1,I0	VAR6070
10 WRITE (6,IOFT) I,(A(I,J),J=1,J0)	VAR6080
11 CONTINUE	VAR6090
RETURN	VAR6100
END	VAR6110
C*****	VAR6120
SUBROUTINE ZERO (A)	VAR6130
C *****	VAR6140
C * SUBROUTINE TO ZERO ARRAYS AT THE START OF PROCESSING FOR EACH LAYER*	VAR6150
C *****	VAR6160
COMMON /MISC/	VAR6170
+ IO, J0	VAR6180
DIMENSION	VAR6190
+ A(I0,J0)	VAR6200
C	VAR6210
C	VAR6220
DO 2 I=1,I0	VAR6230
DO 1 J=1,J0	VAR6240
1 A(I,J)=0.0	VAR6250
2 CONTINUE	VAR6260
RETURN	VAR6270
END	VAR6280

# APPENDIX C.--PROGRAM STEAM

```

C *****
C *                                     PROGRAM STEAM                                *STM0010
C *                                     -----                                *STM0015
C *PROGRAM TO COMPUTE DENSITY OF GROUND WATER AND DENSITY OF PURE WATER*STM0020
C * COMPONENT USING PRELIMINARY STEAM TABLES OF POTTER,R.W., AND D.L. *STM0030
C * BROWN,1977,"THE VOLUMETRIC PROPERTIES OF AQUEOUS SODIUM CHLORIDE *STM0040
C * SOLUTIONS FROM 0-500 DEGREES C. AT PRESSURES UP TO 2000 BARS BASED *STM0050
C * ON A REGRESSION OF AVAILABLE DATA IN THE LITERATURE", GEOLOGICAL *STM0060
C * SURVEY BULLETIN 1421-C. *STM0070
C * PROGRAM ALSO COMPUTES VISCOSITY OF GROUND WATER FROM FIT TO DATA OF*STM0080
C * MATTHEWS,C.S.,AND D.G. RUSSEL,1967,"PRESSURE BUILDUP AND FLOW TESTS*STM0090
C * IN WELLS",APPENDIX G: DALLAS,TEXAS, SOC. OF PETROLEUM ENGS. OF THE *STM0100
C * AM.INST. OF MINING,METALLURGICAL, AND PETROLEUM ENGS. MONOGRAPH #1,*STM0110
C * P. 158. *STM0120
C * ORIGINAL PROGRAM WRITTEN AND DOCUMENTED BY EMANUEL WEISS,1982, *STM0130
C * "A MODEL FOR THE SIMULATION OF VARIABLE-DENSITY GROUND WATER IN *STM0140
C * THREE DIMENSIONS UNDER STEADY-STATE CONDITIONS", U.S.GEOLOGICAL *STM0150
C * SURVEY OPEN-FILE REPORT 82-352. *STM0160
C * THIS VERSION OF PROGRAM MODIFIED BY R.J. MANDLE AND A.L. KONTIS. *STM0170
C * MODIFICATIONS CONSIST OF COMPUTING TEMPERATURE DISTRIBUTION FROM *STM0180
C * ASSUMED GEOTHERMAL GRADIENT AND TO USE BICUBIC INTERPOLATION OF *STM0190
C * POTTER AND BROWN, TABLE 29 TO OBTAIN Ca,Cb,Cc COEFFICIENTS. *STM0200
C * PROCEEDURE TO OBTAIN INTERPOLATION COEFFICIENTS GIVEN IN DAVIS,T.M.*STM0210
C * AND KONTIS,A.L.,1970," SPLINE INTERPOLATION ALGORITHMS FOR TRACK- *STM0220
C * TYPE DATA WITH APPLICATION TO THE COMPUTATION OF MEAN GRAVITY ANO- *STM0230
C * MALIES", U.S. NAVAL OCEANOGRAPHIC OFFICE TECHNICAL REPORT-226. *STM0240
C *                                     INPUT DATA                                *STM0250
C *                                     -----                                *STM0260
C * READ INPUT DATA FROM FORTRAN UNIT NUMBERS AS FOLLOWS: *STM0270
C * CONTROL RECORD (CDIFF,IPRINT,IO,JO) FROM UNIT 5 *STM0280
C * DISSOLVED SOLIDS (TDS) IN GRAMS PER LITER FROM UNIT 10 *STM0290
C * FRESHWATER HEADS (H) IN FEET FROM UNIT 11 *STM0300
C * NODE ELEVATIONS (ELEV) IN FEET FROM UNIT 12 *STM0310
C * POTTER TABLE 29 -Ca BICUBIC COEFFICIENTS (BC) FROM UNIT 13 *STM0320
C * POTTER TABLE 29 -Cb BICUBIC COEFFICIENTS (BC) FROM UNIT 14 *STM0330
C * POTTER TABLE 29 -Cc BICUBIC COEFFICIENTS (BC) FROM UNIT 15 *STM0340
C *                                     OUTPUT DATA                                *STM0350
C *                                     -----                                *STM0360
C * WRITE OUTPUT DATA ON FORTRAN UNIT NUMBERS AS FOLLOWS: *STM0370
C * PROGRAM PRINTOUT ON UNIT 6 *STM0380
C * VISCOSITY (ZMU) IN CENTIPOISE ON UNIT 20 *STM0390
C * GROUND WATER DENSITY (XDEN) IN KG. PER CU.METER ON UNIT 21 *STM0400
C * PURE WATER DENSITY (XDENP) IN KG. PER CU.METER ON UNIT 22 *STM0410
C *                                     PROGRAM INTERNAL ARRAY NAMES *STM0420
C *                                     -----                                *STM0430
C * OTHER ARRAY NAMES USED IN PROGRAM ARE AS FOLLOWS: *STM0440
C * P IS PRESSURE IN BARS *STM0450
C * TEMP IS TEMPERATURE IN DEGREES CELSIUS FOR DENSITY CALCULATIONS *STM0460
C * AND IN DEGREES FARENHEIT FOR VISCOSITY CALCULATIONS *STM0470
C * DEN IS GROUND WATER DENSITY IN GRAMS PER CUBIC CM. *STM0480
C * DENP IS PURE WATER DENSITY IN GRAMS PER CUBIC CM. *STM0490
C * *STM0500

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# APPENDIX C (Continued)

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C * SUBROUTINES CONSIST OF: SUBROUTINE ARRAYS TO READ INPUT ARRAYS AND *STM0510
C * SUBROUTINE BICUBE TO READ BICUBIC COEFFICIENTS AND TO PERFORM *STM0520
C * INTERPOLATION OF POTTER TABLE 29 VALUES. *STM0530
C * DIMENSIONING IS SET BY PARAMETER STATEMENTS IN MAIN PROGRAM AND *STM0540
C * IN SUBROUTINE ARRAYS: NR=NUMBER OF ROWS, NC=NUMBER OF COLUMNS *STM0550
C * NOTE THAT PROGRAM IS WRITTEN TO COMPUTE ONE LAYER AT A TIME. *STM0560
C * IF THERE ARE KO LAYERS, RUN PROGRAM KO TIMES. *STM0565
C *****STM0570
    PARAMETER NR=37,NC=40 STM0580
    DIMENSION H(NR,NC),INFT(2,2),IOFT(9,4),ZMU(NR,NC) STM0590
    DIMENSION TEMP(NR,NC), ELEV(NR,NC), TDS(NR,NC), P(NR,NC) STM0600
    DIMENSION DEN(NR,NC), DENP(NR,NC), XDEN(NR,NC), XDENP(NR,NC) STM0610
    DIMENSION E(3) STM0620
    COMMON/MISC/IO,JO STM0630
C --- E(I) ARE CONSTANTS IN POTTER'S EQ 1 FOR 20 DEGREES CENTIGRADE STM0640
    DATA E/15.782,2.0324,0.0744/ STM0650
    DATA INFT/4H(20F,4H4.0),4H(10F,4H8.0)/ STM0660
    DATA IOFT/'(1H0','I2','2X,2','0F6.','1/(5','X,20','F6.1',')) ',' STM0670
1 ' ','(1H0','I5','14F9','5/(','6X,1','4F9.','5)) ',' STM0680
2 ' ','(1H0','I5','10E1','2.5/','(6X','10E1','2.5)',' ) ',' STM0690
3, '(1H0','I5','10E1','1.3/','(6X','10E1','1.3)',' ) ',' / STM0700
C SPECIFY THE MOLECULAR WEIGHT OF NACL (WT) AND FACTOR (CF) TO STM0710
C CONVERT FEET OF WATER TO BARS (PRESSURE) STM0720
    WT=58.4428 STM0730
    CF=.02987712 STM0740
C---- READ IN CLOSURE CRITERION(CDIFF) FOR MOLALITY ITERATION IN PERCENTSTM0750
C---- ,PRINT CODE(IPRINT), NUMBER OF MODEL ROWS (IO), AND COLUMNS (JO) STM0760
C---- FOR PARTIAL PRINTOUT ON UNIT 6, SET IPRINT=0. STM0770
C FOR FULL PRINTOUT OF CALCULATIONS AT EACH ITERATION SET IPRINT=1 STM0780
    READ (5,250) CDIFF,IPRINT,IO,JO STM0790
    WRITE (6,260) CDIFF,IPRINT STM0800
    WRITE (6,280) IO,JO STM0810
C---- READ IN DISSOLVED SOLIDS (IN GRAMS PER LITER) STM0820
    DO 97 I=1,IO STM0830
    READ(10,98)(TDS(I,J),J=1,JO) STM0840
    98 FORMAT(10F8.3) STM0850
    97 CONTINUE STM0860
C---- READ IN 'FRESHWATER' HEADS (PURE WATER AT STP) IN FEET STM0870
    CALL ARRAYS(H,INFT(1,2),IOFT(1,1),11) STM0880
C---- READ IN NODE ELEVATION (IN FEET RELATIVE TO MSL) STM0890
    CALL ARRAYS(ELEV,INFT(1,2),IOFT(1,4),12) STM0900
    IOM1=IO-1 STM0910
    JOM1=JO-1 STM0920
    DO 110 J=1,JO STM0930
    DO 110 I=1,IO STM0940
    DEN(I,J)=0.0 STM0950
    DENP(I,J)=0.0 STM0960
    110 P(I,J)=0.0 STM0970
C STM0980
C COMPUTE TEMPERATURE AS FUNCTION OF DEPTH FROM GEOTHERMAL GRADIENT STM0990

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# APPENDIX C (Continued)

C	OF 0.00083 C/FT +17.09 OBTAINED BY FITTING STRAIGHT LINE TO DATA	STM1000
C	FROM ILLINOIS BASIN.	STM1010
C		STM1020
	DO 121 I=1,I0	STM1030
	DO 125 J=1,J0	STM1040
	TEMP(I,J)=0.0083*(H(I,J)-ELEV(I,J))+17.09	STM1050
125	CONTINUE	STM1060
	WRITE(6,126)(TEMP(I,J),J=1,J0)	STM1070
126	FORMAT(1X,10F8.2)	STM1080
121	CONTINUE	STM1090
C		STM1100
C	***** START MAJOR LOOP TO COMPUTE MOLALITY AND DENSITIES *****	STM1110
	DO 220 I=2,IOM1	STM1120
	DO 220 J=2,JOM1	STM1130
	P(I,J)=CF*(H(I,J)-ELEV(I,J))	STM1140
	XMOL=TDS(I,J)*1000/((1000-TDS(I,J))*WT)	STM1150
	N=0	STM1160
130	CONTINUE	STM1170
C	--- DEMO IS DENSITY OF PURE WATER AT 20 DEG.C. AND 1BAR	STM1180
C	INTERPOLATED FROM TABLE 29 OF POTTER AND BROWN.	STM1185
	DENO = 0.9976054	STM1190
	DNM=1000+(E(1)*XMOL+E(2)*(SQRT(XMOL)**3)+E(3)*(XMOL**2))*DENO	STM1200
	DEN1=(1000*DENO+WT*XMOL*DENO)/DNM	STM1210
	N=N+1	STM1220
	XMOL1=TDS(I,J)*1000/((DEN1*1000-TDS(I,J))*WT)	STM1230
	CRTN=ABS((XMOL-XMOL1)*100/XMOL)	STM1240
	XMOL=XMOL1	STM1250
	IF (N.EQ.1.AND.IPRINT.EQ.1) WRITE (6,320)	STM1260
	IF (PRINT.EQ.1) WRITE (6,330) I,J,N,XMOL1,DEN1,TDS(I,J),CRTN	STM1270
	IF(CRTN.GT.CDIFF.AND.N.LT.10) GO TO 130	STM1280
	IF(N.LE.9) GO TO 140	STM1290
	WRITE (6,340) I,J,N,DEN1,XMOL1,CRTN	STM1300
140	CONTINUE	STM1310
C	FOR MOLALITY AND PRESSURE OF NODE I,J CALL SUBROUTINE BICUBE THREE	STM1320
C	TIMES TO OBTAIN POTTER TABLE 29 VALUES (ACOE= Ca, BCOEF= Cb,	STM1330
C	CCOE= Cc) AND REWIND FILES FOR NEXT NODE COMPUTATION.	STM1340
	LIN=13	STM1350
	PP=P(I,J)	STM1360
	CALL BICUBE(XMOL,PP,ACOE,LIN)	STM1370
	REWIND 13	STM1380
	LIN=14	STM1390
	CALL BICUBE(XMOL,PP,BCOE,LIN)	STM1400
	REWIND 14	STM1410
	LIN=15	STM1420
	CALL BICUBE(XMOL,PP,CCOE,LIN)	STM1430
	REWIND 15	STM1440
C		STM1450
C	COMPUTE GROUND WATER DENSITY	STM1460
C		STM1470
	RHO=ACOE-(1.E-4)*BCOE*TEMP(I,J)-(1.E-6)*CCOE*(TEMP(I,J)**2.)	STM1480
180	CONTINUE	STM1490
C		STM1500



# APPENDIX C (Continued)

C	COMPUTE PURE WATER DENSITY	STM1510
C		STM1520
	DEN(I,J)=RHO	STM1530
	DENP(I,J)=DEN(I,J)*(1-XMOL*WT/(XMOL*WT+1000))	STM1540
220	CONTINUE	STM1550
C		STM1560
C	CONVERT DENSITIES TO KILOGRAMS PER CUBIC METER	STM1570
C		STM1580
	DO 230 J=1,J0	STM1590
	DO 230 I=1,I0	STM1600
	XDEN(I,J)=(DEN(I,J)*1000)	STM1610
230	XDENP(I,J)=(DENP(I,J)*1000)	STM1620
C		STM1630
C	CALCULATE VISCOSITY FROM GRAPH IN MATTHEWS AND RUSSELL,1967	STM1640
C		STM1650
	DO 232 I=1,I0	STM1660
	DO 231 J=1,J0	STM1670
	TEMP(I,J)=TEMP(I,J)*9/5+32	STM1680
	XMU=38.3432/SQRT(TEMP(I,J))-14.623/SQRT(SQRT(TEMP(I,J)))	STM1690
	YMU=XMU+1.481	STM1700
	TDSMU=1.+(TDS(I,J)/300)	STM1710
	ZMU(I,J)=YMU*TDSMU	STM1720
231	CONTINUE	STM1730
C		STM1740
C	WRITE OUTPUT ARRAYS	STM1750
C		STM1760
	WRITE(6,391)	STM1770
	WRITE(20,392)(ZMU(I,J),J=1,J0)	STM1780
391	FORMAT(1X,'VISCOSITY DATA')	STM1790
392	FORMAT(10F8.3)	STM1800
232	CONTINUE	STM1810
	WRITE (6,360)	STM1820
	DO 393 I=1,I0	STM1830
	WRITE (6,390) (XDEN(I,J),J=1,J0)	STM1840
393	CONTINUE	STM1850
	WRITE (6,370)	STM1860
	DO 394 I=1,I0	STM1870
	WRITE (6,390) (XDENP(I,J),J=1,J0)	STM1880
	WRITE (21,390) (XDEN(I,J),J=1,J0)	STM1890
	WRITE (22,390) (XDENP(I,J),J=1,J0)	STM1900
394	CONTINUE	STM1910
	WRITE(6,789)	STM1920
789	FORMAT(1H0,' PRESSURE ARRAY')	STM1930
	DO 787 I=1,I0	STM1940
787	WRITE(6,390) (P(I,J),J=1,J0)	STM1950
	STOP	STM1960
C		STM1970
250	FORMAT (F10.6,3I10)	STM1980
260	FORMAT ('1',T35,'ITERATION CLOSURE CRITERION FOR DENSITY(IN %) =',	STM1990
	1F10.6,20X,'PRINT ='F4.1)	STM2000
270	FORMAT (3I10)	STM2010
280	FORMAT (T45,'I0 J0 ',3I3)	STM2020

# APPENDIX C (Continued)

```

320 FORMAT (/ ,T25,'XMOL1',6X,'DEN1',7X,'TDS',7X,'CRTN')          STM2030
330 FORMAT (T2,3I5,4F10.5)                                         STM2040
340 FORMAT(T2,'GT 10 ITERATIONS',3I5,'DN1=',F10.6,'XMOL=',F10.6,'CRTN=STM2050
    1',F8.4)                                                         STM2060
360 FORMAT ('1',T22,'GROUND WATER DENSITY (G/L)',/)              STM2070
370 FORMAT ('1',T22,'PURE WATER CONCENTRATION (G/L)',/)          STM2080
390 FORMAT(10F8.1)                                                 STM2090
    END                                                             STM2100
C *****                                                         STM2110
    SUBROUTINE ARRAYS(A,INFT,IOFT,LIN)                             STM2120
C *****                                                         STM2130
    PARAMETER NR=37,NC=40                                         STM2140
    DIMENSION A(NR,NC),INFT(2,3),IOFT(9,4)                       STM2150
    COMMON /MISC/ IO,J0                                           STM2160
    DO 50 I=1,IO                                                  STM2170
    READ (LIN,INFT) (A(I,J),J=1,J0)                               STM2180
50 CONTINUE                                                         STM2190
    DO 70 I=1,IO                                                  STM2200
70 WRITE (6,IOFT) I,(A(I,J),J=1,J0)                               STM2210
80 CONTINUE                                                         STM2220
    RETURN                                                         STM2230
    END                                                             STM2240
C *****                                                         STM2330
    SUBROUTINE BICUBE(XMOL,P,COEFF,LIN)                             STM2340
C *****                                                         STM2350
C * SUBROUTINE TO OBTAIN VALUES OF Ca,Cb,Cc COEFFICIENTS OF TABLE 29, *STM2360
C * (POTTER AND BROWN, GEOLOGICAL SURVEY BULLETIN 1421-C) FOR GIVEN *STM2370
C * VALUES OF MOLALITY AND PRESSURE. Ca,Cb,Cc VALUES COMPUTED FROM *STM2380
C * BICUBIC USING COEFFICIENTS GENERATED FROM TABULATED VALUES OF *STM2390
C * TABLE 29. ESTIMATES OF Ca,Cb,Cc FOR MOLALITY OF 0, AND PRESSURES UP*STM2400
C * TO 400 BARS PROVIDED BY R.J.SUN (WRITTEN COMMUNICATION,1981). *STM2410
C * *                                                             *STM2420
C * FOR EACH SET OF TABLE VALUES (TABLE 2 OF THIS REPORT), *STM2430
C * THERE ARE 12x4 CELLS. FOR EACH CELL, THERE ARE 16 BICUBIC COEF- *STM2440
C * FICIENTS SO THAT THE BICUBIC COEFFICIENT ARRAYS ARE OF SIZE 48x16. *STM2450
C * IN SUBROUTINE BICUBE THE 16 COEFFICIENTS FOR EACH OF THE 4 COLUMN *STM2460
C * CELLS OF A PARTICULAR ROW ARE READ AND PLACED IN ARRAY BC OF SIZE *STM2470
C * 4x4x4. IF THE ROW OF THE COMPUTED MOLALITY CORRESPONDS TO THE ROW *STM2480
C * OF THE BICUBIC COEFFICIENTS THEN THE 16 COEFFICIENTS CORRESPONDING *STM2490
C * TO THE COMPUTED PRESSURE CELL ARE PLACED IN ARRAY C OF SIZE 4x4. *STM2500
C * THESE COEFFICIENTS ARE THEN USED IN THE BICUBIC POLYNOMIAL TOGETHER*STM2510
C * WITH THE COMPUTED MOLALITY AND PRESSURE TO OBTAIN THE INTERPOLATED *STM2520
C * TABLE VALUE. UPON RETURNING TO THE MAIN PROGRAM THE BICUBIC COEF- *STM2530
C * FICIENT FILE IN QUESTION IS REWOUND. *STM2540
C *****                                                         STM2550
    DOUBLE PRECISION C                                           STM2560
    DIMENSION BC(4,4,4),C(4,4)                                     STM2570
    ICOL=12                                                         STM2580
    JCOL=4                                                           STM2590
    YMOLE=XMOL*2.0                                                 STM2600
    II=IFIX(YMOLE)+1                                              STM2610

```

# APPENDIX C (Continued)

JJ=IFIX(P/100.)+1	STM2620
DO 69 I=1,ICOL	STM2630
DO 70 J=1,JCOL	STM2640
70 READ(LIN,60)((BC(M,N,J),M=1,4),N=1,4)	STM2650
60 FORMAT(10F8.4)	STM2660
IF(II.NE.1) GO TO 69	STM2670
DO 79 N=1,4	STM2680
DO 78 M=1,4	STM2690
78 C(M,N)=BC(M,N,JJ)	STM2700
79 CONTINUE	STM2710
GO TO 300	STM2720
69 CONTINUE	STM2730
300 CONTINUE	STM2740
IP=(IFIX(P/100.))*100	STM2750
X=(P-IP)/100.	STM2760
Y=YMOLE-II+1	STM2770
X2=X*X	STM2780
X3=X2*X	STM2790
Y2=Y*Y	STM2800
Y3=Y2*Y	STM2810
COEFF= C(1,1)+C(1,2)*X+C(1,3)*X2+C(1,4)*X3+Y*(C(2,1)+C(2,2)*X+	STM2820
1C(2,3)*X2+C(2,4)*X3)+Y2*(C(3,1)+C(3,2)*X+C(3,3)*X2+C(3,4)*X3)+	STM2830
2Y3*(C(4,1)+C(4,2)*X+C(4,3)*X2+C(4,4)*X3)	STM2840
RETURN	STM2850
END	STM2860

# APPENDIX D.--WATER-DENSITY MODIFICATIONS TO

## TRESCOTT-LARSON MODEL PROGRAM

```

*****
*
*           TRESCOTT MAIN PROGRAM--MODIFICATIONS
*           FOR VARIABLE DENSITY
*
*           _____
*
*****
      DIMENSION Y(120000),L(30),HEADNG(33),NAME(54),INFT(2,4),      IOFT(MANO100
3T,4HOM E,4HLEVA,4HTION,2*4H      ,4H      R,4HECHA,4HRGE ,4HRATE,2*4H      MANO240
4      ,4H TR ,4HCOEF,4HFICI,4HENT ,2*4H      ,4H TC ,4HCOEF,4HFICI,4HENTMANO241
5 /      MANO242
      DATA INFT/4H(20F,4H4.0),4H(10F,4H8.0),4H(8F1,4H0.0),4H(8E1,4H0.4)/MANO250
      ISUM=ISUM+ISIZ      MANI050
C PLACE STARTING HEADS AND OUTPUT FROM VARDEN PROGRAM CONSISTING OF:      MANI381
C STORAGE COEFF.,TRANSMISSIVITY,VERTICAL LEAKANCE(TK),VARIABLE DENSITY      MANI382
C SOURCE TERM(PST),TR CONDUCTANCE,AND TC CONDUCTANCE IN INPUT FILE LIN.      MANI383
      LIN=11      MANI384
      80 CALL ARRAY(Y(LOC),INFT(1,2),IOFT(1,1),NAME(1),IRN,DUM,LIN)      MANI410
      90 CALL ARRAY(Y(LOC),INFT(1,1),IOFT(1,2),NAME(7),IRN,DUM,LIN)      MANI440
      CALL ARRAY(Y(LOC),INFT(1,4),IOFT(1,2),NAME(13),IRN,DUM,LIN)      MANI500
      110 CALL ARRAY(Y(LOC),INFT(1,4),IOFT(1,3),NAME(19),IRN,DUM,LIN)      MANI580
C FOR VARIABLE DENSITY REPLACE TRESCOTT STATEMENTS      MANI620-MANI640      MANI611
C WITH THE FOLLOWING STATEMENTS.      MANI612
C FOR VARIABLE DENSITY SET IQRE OPTION ON OPTIONS RECORD.      MANI613
      130 IF(IQRE.NE.ICHK(7)) GO TO 141      MANI614
C READ IN VARIABLE SOURCE TERM (PST)      MANI615
      DO 142 K=1,K0      MANI616
      LOC=L(25)+(K-1)*NIJ      MANI617
      142 CALL ARRAY(Y(LOC),INFT(1,4),IOFT(1,3),NAME(37),IRN,DUM,LIN)      MANI618
C      MANI619
C READ IN TR COEFFICIENTS      MANI620
C      MANI621
      DO 143 K=1,K0      MANI622
      LOC=L(6)+(K-1)*NIJ      MANI623
      143 CALL ARRAY(Y(LOC),INFT(1,4),IOFT(1,3),NAME(43),IRN,DUM,LIN)      MANI624
C      MANI625
C READ IN TC COEFFICIENTS      MANI626
C      MANI627
      DO 144 K=1,K0      MANI628
      LOC=L(7)+(K-1)*NIJ      MANI629
      144 CALL ARRAY(Y(LOC),INFT(1,4),IOFT(1,3),NAME(49),IRN,DUM,LIN)      MANI631
      141 CONTINUE      MANI632
      CALL MDAT      MANI640
*****
*
*           TRESCOTT SUBROUTINE DATAI--MODIFICATIONS
*           FOR VARIABLE DENSITY
*
*****

```

# APPENDIX D (Continued)

```

C MX=NUMBER OF MODEL ROWS,NX=NUMBER OF MODEL COLUMNS.          DAT 71
  PARAMETER MX=37,NX=40                                          DAT 73
  3TOM(IP,JP),QRE(IO,J0,K0), TF(3), A(MX,NX), IN(6), IOFT(9), INFT(2)DAT 140
  ENTRY ARRAY(A,INFT,IOFT,IN,IRN,TF,LIN)                        DAT 710
C THE ORDER OF VARDEN PROGRAM OUTPUT ARRAYS IS FROM BOTTOM ROW TO TOP DAT 815
C ROW, SO REVERSE ROW ORDER OF READ STATEMENT TO CONFORM TO TRESCOTT DAT 816
C CONVENTION.                                                    DAT 817
  DO 110 I=IO,1,-1                                              DAT 820
  READ (LIN,INFT) (A(I,J),J=1,J0)                               DAT 830
*****
*
*       TRESCOTT SUBROUTINE CHECKI--MODIFICATIONS                *
*       FOR VARIABLE DENSITY                                     *
*
*****
  3W(NCH), QRE(IO,J0,K0)                                         CHK 130
C NOTE THAT FOR VARIABLE DENSITY OPTION, COMPUTATION OF FLOW FOR NODES
C WITH RELATIVE DENSITY OTHER THAN UNITY REQUIRES ADDITION OF APPRO-
C PRIATE SOURCE TERM (ALPHA,BETA,OR GAMMA FROM PROGRAM VARDEN).
C SUBROUTINE CHECKI WAS NOT MODIFIED TO COMPUTE FLOW WITH VARIABLE DEN-
C SITY SO THAT THE FLOW BETWEEN CONSTANT HEAD NODES AND ADJACENT ACTIVE
C NODES MAY NOT BE CORRECT. BECAUSE OF THIS THE PORTION OF THE MODEL
C MASS BALANCE OUTPUT ATTRIBUTED TO FLOW FROM/TO CONSTANT HEAD NODES
C AND FLOW FROM/TO THE TOP LAYER WILL BE APPROXIMATE. ACTUAL FLOW MAY
C BE COMPUTED WITH A POST-MODEL PROCESSOR BY SAVING MODEL HEADS AND
C USING THE TR,TC,TK CONDUCTANCES AND ALPHA,BETA,GAMMA SOURCE TERMS
C FROM PROGRAM VARDEN.
C ADD TOTAL SOURCE TERM TO PUMP OR CFLUX                         CHK1019
C (K.EQ.K0.AND.IQRE.EQ.ICHK(7)) QREFLX=QREFLX+QRE(I,J,K)*AREA   CHK1020
  IF(QRE(I,J,K)) 155,158,157                                     CHK1021
155 PUMP=PUMP+QRE(I,J,K)                                         CHK1022
  GO TO 158                                                       CHK1023
157 CFLUX=CFLUX+QRE(I,J,K)                                       CHK1024
158 CONTINUE                                                     CHK1025
*****
*
*       TRESCOTT SUBROUTINE COEFF MODIFICATIONS                *
*       FOR VARIABLE DENSITY                                     *
*
*****
  3TOM(IP,JP),QRE(IO,J0,K0)                                         COF 130
C BYPASS TR AND TC CONDUCTANCE CALCULATIONS IF VARIABLE DENSITY OPTION COF 512
C IS USED.                                                         COF 513
  IF(IQRE.EQ.ICHK(7)) GO TO 333                                    COF 514
333 CONTINUE                                                     COF 741

```

# APPENDIX E.--MULTIAQUIFER-WELL MODIFICATIONS TO

## TRESCOTT-LARSON MODEL PROGRAM

```

*****
*
*           TRESCOTT MAIN PROGRAM--MODIFICATIONS
*           FOR MULTIAQUIFER WELL
*
*
*****
PARAMETER MW=365,MJ=8,MQ=4                                MAN0091
2H, IDK1, IDK2, IWATER, IQRE, IP, JP, IQ, JQ, IK, JK, K5, IPU1, IPU2, ITK, IEQN MAN0170
3, IOPEN, IALOCA                                           MAN0175
COMMON/WIS/IL(MW), JL(MW), NWL(MW), KLL(MQ, MW, MJ), RWL(MW), QWL(MJ, MAN0201
1MW), C2(MJ, MW), C1(MQ, MW), NU(MJ, MW), NR(MQ, MW), PI2   MAN0202
DATA PI2/6.28318/                                           MAN0291
1, IEQN, IOPEN, IALOCA                                     MAN0395
1, IEQN, IOPEN, IALOCA                                     MAN0405
*****
*
*           TRESCOTT SUBROUTINE DATAI--MODIFICATIONS
*           FOR MULTIAQUIFER WELL
*
*
*****
C MW=NUMBER OF NODE SETS                                     *
C MJ=MAXIMUM NUMBER OF MULTIAQUIFER CATEGORIES.             DAT 72
C FOR NORTHERN MIDWEST TOP LAYER IS CONSTANT HEAD SO THAT MQ=K0-1, DAT 73
C OTHERWISE MQ IS NUMBER OF ACTIVE LAYERS.                  DAT 74
PARAMETER MW=365,MJ=8,MQ=4                                  DAT 73
COMMON/WIS/IL(MW), JL(MW), NWL(MW), KLL(MQ, MW, MJ), RWL(MW), QWL(MJ, DAT 241
1MW), C2(MJ, MW), C1(MQ, MW), NU(MJ, MW), NR(MQ, MW), PI2   DAT 242
C FOR MULTIAQUIFER WELL MODIFACTIONS REPLACE TRESCOTT STATEMENTS
C DAT1740--DAT1820 OF ORIGINAL PROGRAM WITH THE FOLLOWING.
C --- READ AND WRITE PUMPING RATES ---                      DAT1740
DO 245 K=1, K0                                              DAT1741
DO 245 I=1, IO                                              DAT1742
DO 245 J=1, JO                                              DAT1743
245 WELL(I, J, K)=0.0                                       DAT1744
C***** LOOP TO 333 IF OPEN WELL *****                   DAT1745
IF(IOPEN.EQ.ICHK(12)) GO TO 333                             DAT1746
WRITE(6,410) NWEL                                           DAT1747
IF(NWEL.EQ.0) GO TO 260                                     DAT1748
DO 270 II=1, NWEL                                           DAT1749
READ(5,330) K, I, J, WELL(I, J, K)                         DAT1750
270 WELL(I, J, K)=WELL(I, J, K)/(DELX(J)*DELY(I))          DAT1751
GO TO 260                                                    DAT1758
333 CONTINUE                                                DAT1759
KK=MQ                                                        DAT1760
DO 250 I=1, NWEL                                           DAT1761
READ(11,500) IL(I), JL(I), JA, ((KLL(K, I, J), K=1, KK), QWL(J, I), J=1, JA) DAT1762
C WRITE(6,501) IL(I), JL(I), JA, ((KLL(K, I, J), K=1, KK), QWL(J, I), J=1, JA) DAT1763
501 FORMAT(1X,2I2,4X,I2,8(4I1,F6.3))                      DAT1764
NWL(I)=JA                                                    DAT1765

```

# APPENDIX E (Continued)

C	COMPUTE ESTIMATE OF NO. OF WELLS PER MULTIAQUIFER COMBINATION	DAT1766
	AVPUMP=0.1	DAT1767
	DO 692 JJ=1,JA	DAT1768
	IF(QWL(JJ,I).EQ.0.0) NU(JJ,I)=1	DAT1769
	IF(QWL(JJ,I).EQ.0.0) GO TO 692	DAT1770
	NU(JJ,I)=QWL(JJ,I)/AVPUMP	DAT1771
	IF(NU(JJ,I).EQ.0) NU(JJ,I)=1	DAT1772
	692 CONTINUE	DAT1773
C		DAT1774
C	COMPUTE NO. OF WELLS TAPPING LAYER K(NR(K,I))	DAT1775
C	I.E. NR(K,I) CONTAINS THE TOTAL NUMBER OF WELLS OPEN TO LAYER K	DAT1776
C	OF NODE SET I.	DAT1777
C		DAT1778
	DO 688 K=1,KK	DAT1779
	NRSUM=0	DAT1780
	DO 687 JJ=1,JA	DAT1781
	687 IF(KLL(K,I,JJ).NE.0) NRSUM=NRSUM+NU(JJ,I)	DAT1782
	NR(K,I)=NRSUM	DAT1783
	688 CONTINUE	DAT1784
C		DAT1785
C	FOR THE NORTHERN MIDWEST THE PUMPING RATE FILE HAS POSITIVE VALUES	DAT1786
C	FOR DISCHARGE. HENCE SIGN OF PUMPING RATES IS CHANGED ON INPUT.	DAT1787
C	IN ADDITION, BECAUSE VARIABLE DENSITY WAS SIMULATED PUMPING RATES	DAT1788a
C	ARE MULTIPLIED BY DENSITY OF PURE WATER.	DAT1788b
	DO 7356 J=1,JA	DAT1789
	7356 QWL(J,I)=-1.5473*QWL(J,I)*1.94	DAT1790
	DX=DELX(JL(I))	DAT1791
	DY=DELY(IL(I))	DAT1792
C	IF VARIABLE DENSITY OPTION IS USED THEN DX=DY=1 FOR ALL NODES.	DAT1793
C	TO COMPUTE EFFECTIVE RADIUS (REL) THE AREA OF THE BLOCK IS NEEDED.	DAT1794
C	BECAUSE A UNIFORM NODE SPACING WAS USED IN NORTHERN MIDWEST(16X16	DAT1795
C	MILES) AREA TERM BELOW IS EXPLICITLY DEFINED. IF A NON-UNIFORM	DAT1796a
C	SPACING IS USED AND VARIABLE DENSITY IS ALSO BEING SIMULATED,PRO-	DAT1796b
C	GRAM MUST BE MODIFIED TO ALLOW COMPUTATION OF AREA.	DAT1796c
C	IN GENERAL, AREA=MIN(DX,DY)**2.	DAT1796d
	AREA=84480.*84480.	DAT1797
	RWL(I)=0.5	DAT1798
C	COMPUTE EFFECTIVE RADIUS FOR EACH LAYER USING NR(K,I)	DAT1799
	DO 686 K=1,KK	DAT1800
	IF(NR(K,I).EQ.0) GO TO 686	DAT1801
	REL=SQRT(0.04322*AREA/NR(K,I))	DAT1802
	C1(K,I)=1.0/(ALOG(REL/RWL(I)))	DAT1803
	686 CONTINUE	DAT1804
	DO 620 J=1,JA	DAT1805
	C2(J,I)=0.0	DAT1806
	KK=MQ	DAT1807
	DO 520 K=1,KK	DAT1808
	KL=KLL(K,I,J)	DAT1809
	IF(KL.EQ.0) GO TO 520	DAT1810
	C2(J,I)=C2(J,I)+T(IL(I),JL(I),KL)*C1(KL,I)	DAT1811

# APPENDIX E (Continued)

520	CONTINUE	DAT1812
	C2(J,I)=1.0/C2(J,I)	DAT1813
620	CONTINUE	DAT1814
	DO 630 K=1, KK	DAT1815
	SUMZ=0.0	DAT1816
	DO 530 J=1, JA	DAT1817
	KL=KLL(K,I,J)	DAT1818
	IF(KL.EQ.0) GO TO 530	DAT1819
	SUMZ=SUMZ+QWL(J,I)*C2(J,I)	DAT1820
530	CONTINUE	DAT1821
	WELL(IL(I),JL(I),K) = T(IL(I),JL(I),K)*SUMZ*C1(K,I)	DAT1822
	WELL(IL(I),JL(I),K)= WELL(IL(I),JL(I),K)/(DX*DY)	DAT1823
630	CONTINUE	DAT1824
250	CONTINUE	DAT1825
500	FORMAT(1X,2I2,I6,8(4I1,F6.0))	DAT1826
510	FORMAT(8F10.0)	DAT1827
260	RETURN	DAT1828
*****		
*		*
*	TRESCOTT SUBROUTINE CHECKI--MODIFICATIONS	*
*	FOR MULTIAQUIFER WELL	*
*		*
*****		
	PARAMETER MW=365,MJ=8,MQ=4	CHK 91
	DIMENSION HW(MJ),HZ(MJ)	CHK 92
	3,IOPEN,IALOCA	CHK 171
	COMMON/WIS/ IL(MW),JL(MW),NWL(MW),KLL(MQ,MW,MJ),RWL(MW),QWL(MJ,	CHK 201
	1MW),C2(MJ,MW),C1(MQ,MW),NU(MJ,MW),NR(MQ,MW),PI2	CHK 202
	IF(IOPEN.NE. ICHK(12)) GO TO 333	CHK1541
	WRITE(6,1111)	CHK1543
	DO 300 I=1,NWEL	CHK1544
	J=JL(I)	CHK1545
	II=IL(I)	CHK1546
C	COMPUTE HEAD IN WELL FOR EACH OF THE "JA" AQUIFERS	CHK1547
	JA=NWL(I)	CHK1548
	DO 310 JJ=1,JA	CHK1549
	SUMZ=0.0	CHK1550
	KK=MQ	CHK1550
	DO 320 L=1, KK	CHK1552
	K=KLL(L,I,JJ)	CHK1553
	IF(K.EQ.0) GO TO 320	CHK1554
	SUMZ=SUMZ+T(II,J,K)*PHI(II,J,K)*C1(K,I)	CHK1555
320	CONTINUE	CHK1556
C	HW IS THE COMPOSITE HEAD WITH PUMPING FOR EACH CATEGORY WELL.	CHK1558
C	HZ IS THE COMPOSITE HEAD WITH NO PUMPING FOR EACH CATEGORY WELL.	CHK1559
C		CHK1561
	HW(JJ)=SUMZ*C2(JJ,I)+QWL(JJ,I)*C2(JJ,I)/(NU(JJ,I)*PI2)	CHK1562
	HZ(JJ)=SUMZ*C2(JJ,I)	CHK1563
311	CONTINUE	CHK1564
310	CONTINUE	CHK1565



# APPENDIX E (Continued)

```

C      WRITE(6,1118)I,II,J,((KLL(L,I,JJ),L=1,KK),QWL(JJ,I),JJ=1,JA)      CHK1566
1118  FORMAT(1X,2(' '), ' NODE SET',I4, ' MODEL NODE(' ',I2,',',',I2,')',2(' '*
1),7(1X,4I1,F8.3)/3(1X,4I1,F8.3))      CHK1568
C      WRITE(6,1117)(HW(JJ),HZ(JJ),JJ=1,JA)      CHK1569
1117  FORMAT(36X,7(F7.0,F6.0)/1(F7.0,F6.0))      CHK1570
C      WRITE COMPOSITE HEAD ON UNIT 13      CHK1570a
      WRITE(13,1115) II,J,JA,((KLL(L,I,JJ),L=1,KK),HZ(JJ),JJ=1,JA)      CHK1571
1115  FORMAT(2I2,1X,I2,1X,8(4I1,F6.0))      CHK1572
1113  FORMAT(3X,'PUMPING RATE (' ',I2,',',',I2,')',10F8.4)      CHK1573
1112  FORMAT(3X,'HEAD IN WELL (' ',I2,',',',I2,')',10F8.2)      CHK1574
      IF(IALOCA.NE.ICHK(13)) GO TO 300      CHK1575
C      BYPASS ALOCAT EXCEPT FOR LAST PUMPING PERIOD      CHK1576
      IF(KP.LT.NPER) GO TO 300      CHK1577
      CALL ALOCAT(PHI,OLD,T,S,TK,DELX,DELY,DELZ,IO,JO,KO,K5,HW,Q,NS1,I,      CHK1578
111,J,TC,TR,QRE)      CHK1579
300  CONTINUE      CHK1580
333  CONTINUE      CHK1581
1111  FORMAT(1H0,'CALCULATED "WELL HEADS" FOR GIVEN PUMPING RATES')      CHK1582
*****
*
*      TRESCOTT SUBROUTINE PRNTA--MODIFICATIONS      *
*      FOR MULTIAQUIFER-WELL      *
*
*****
      3,IOPEN,IALOCA      PRN 151
      1UN2','ITKR','EQN3','OPEN','ALOC'/      BLK 140

```

## APPENDIX F.--SUBROUTINES SOLVE AND BAND

[illegible]

# APPENDIX F (Continued)

C	*****	SP30260
	ENTRY ITER	SP30270
C	*****	SP30280
C	ENTRY POINT ITER ENTERED ONCE PER SIMULATION	0000000
	NIJ=I0*J0	SP30290
	NIJK=NIJ*K0	SP30300
	DO 5 I=1,NIJK	SP30310
	DH(I)=0.	SP30320
5	NP(I)=0	SP30330
C	SET UP PROCEEDURE TO NUMBER EQUATIONS	0000000
	DO 6 K=1,K0	SP30340
6	KPOS(K)=(K-1)*NIJ	SP30350
	DO 7 J=1,J0	SP30360
7	JPOS(J)=(J-1)*I0	SP30370
	NB(1)=0	SP30380
	MBW=K0+1	SP30390
	NEQ=0	SP30400
	ICH(I0)=0	SP30410
C	IDENTIFY ACTIVE MODEL NODES, IDENTIFY SLICES WITH CONSTANT HEADS	0000000
C	(ARRAY ICH),AND COUNT THE TOTAL NUMBER OF EQUATIONS FOR ALL	0000000
C	SLICES(NEQ).	0000000
	DO 11 I=2,I1	SP30420
	ICH(I)=0	SP30430
	NN=0	SP30440
	DO 10 J=2,J1	SP30450
	DO 10 K=1,K0	SP30460
	N=I+JPOS(J)+KPOS(K)	SP30470
	IF(S(N).LT.0.) ICH(I)=1	SP30480
	IF(T(N).EQ.0..OR.S(N).LT.0.) GO TO 10	SP30490
	NN=NN+1	SP30500
	NP(N)=NN	SP30510
	NEQ=NEQ+1	SP30520
10	CONTINUE	SP30530
11	NB(I)=NEQ	SP30540
	NEQBW=NEQ*MBW	SP30550
	PRINT 12,NEQBW	SP30560
12	FORMAT(49H0***** WARNING**ARRAY A MUST BE DIMENSIONED >,I10)	SP30570
C	READ ACCELERATION PARAMETER WO.	0000000
	READ 13,W0	SP30580
13	FORMAT(F10.0)	SP30590
	PRINT 14,W0,LENGTH	SP30600
14	FORMAT(28H ACCELERATION PARAMETER, W0=,F10.3,/,	SP30610
	\$ 49H NUMBER OF ITERATIONS BETWEEN WATT'S CORRECTION =,I5)	SP30620
	RETURN	SP30630
	CAA	00000000
C	*****	SP30640
	ENTRY NEW ITA	SP30650
C	*****	SP30660
C	ENTRY POINT NEW ITA ENTERED EACH TIME STEP.	0000000
C	---COMPUTE TRANSMISSIVITY AND T COEFFICIENTS FOR UPPER	SP30670
C	HYDROLOGIC UNIT WHEN IT IS UNCONFINED---	SP30680

# APPENDIX F (Continued)

```

        IF (IWATER.NE.ICHK(6)) GO TO 35                      SP30690
        CALL TRANS(0)                                         SP30700
35  CONTINUE                                                  SP30710
        IR=0                                                  SP30720
        ID=1-MBW                                              SP30730
        DO 36 I=1,NEQBW                                       SP30740
36  A(I)=0.                                                    SP30750
        RESO=0.0                                              ***0760
C  CONSTRUCT COEFFICIENTS AND LOAD INTO A AND RHS          SP30770
C *****0000000
C ***** CALL SUBROUTINE MULT FOR MULTIAQUIFER OPTION ***** ***0000
        IF(IOPEN.NE.ICHK(12)) GO TO 333                      ***0771
        CALL MULT(NWEL,A,NB,MBW,JPOS,KPOS,T,NP,KO)           ***0772
        IWELL=1                                               ***0773
333  CONTINUE                                                ***0774
C *****0000000
CBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBBB00000000
C  START SLICES TO DEVELOP THE B,D,E,F,H,SU,Z COEFFICIENTS AND LOAD 0000000
C  THEM INTO 'A' ARRAY. ALSO DEVELOP THE RHS PART OF THE RIGHT HAND 0000000
C  SIDE OF THE EQUATIONS.                                     0000000
        DO 60 I=2,I1                                          SP30780
        DO 47 J=2,J1                                          SP30790
        IFLAG=0                                               ***0791
        DO 48 K=1,K0                                          SP30800
        N=I+JPOS(J)+KPOS(K)                                   SP30810
        IF(NP(N).EQ.0) GO TO 48                               SP30820
        NIA=N+1                                               SP30830
        NIB=N-1                                               SP30840
        NJA=N+IO                                              SP30850
        NJB=N-IO                                              SP30860
        NKA=N+NIJ                                             SP30870
        NKB=N-NIJ                                             SP30880
        VX=DELX(J)                                           SP30890
        VY=DELY(I)                                           SP30900
        VZ=VX*VY                                              SP30910
        VQ=VZ                                                 SP30920
        IF(IEQN.NE.ICHK(11)) GO TO 123                       SP30930
        DZ=DELZ(K)                                           SP30940
        VX=VX*DZ                                              SP30950
        VY=VY*DZ                                              SP30960
        VQ=VQ*DZ                                              SP30970
123  D=TR(NJB)*VY                                             SP30971
        F=TR(N)*VY                                             SP30972
        B=TC(NIB)*VX                                           SP30973
        H=TC(N)*VX                                             SP30974
        SU=0.D0                                               SP30980
        Z=0.D0                                               SP30990
        SUPH=0.0                                              SP31000
        ZPH=0.0                                              SP31010
        IF(K.EQ.1) GO TO 124                                  SP31020
        IF(T(N).EQ.0.0) TK(NKB)=0.0                          ***1030
        Z=TK(NKB)*VZ                                           SP31031

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# APPENDIX F (Continued)

```

      ZPH=Z*PHI(NKB)
124 IF(K.EQ.K0) GO TO 125
      IF(T(N).EQ.0.0) TK(N)=0.0
      SU=TK(N)*VZ
      SUPH=SU*PHI(NKA)
125 RHO=S(N)/DELT
      QR=0.
C *****
C *          VARIABLE-DENSITY MODIFICATIONS
C *          EXPAND NUMBERING ON QRE TERM TO INCLUDE ALL LAYERS
C *****
C      IF(K.NE.K0) GO TO 42
      IF(IQRE.EQ.ICHK(7)) QR=QRE(I+JPOS(J)+KPOS(K))
      42 CONTINUE
C *****
      E=-B-D-F-H-SU-Z-RHO*VQ
C      RES=RHS IS RIGHT HAND SIDE OF EQUATION FOR TIME STEP AND USES
C      HEADS FROM PREVIOUS TIME STEP.
      RES=-B*PHI(NIB)-D*PHI(NJB)-E*PHI(N)-F*PHI(NJA)-H*PHI(NIA)-SUPH-ZPH
1+(-WELL(N) +(-RHO*OLD(N)-QR))*VQ
      IR=IR+1
      ID=ID+MBW
      NN=NP(N)
      A(ID)=E+A(ID)
      IF(K.EQ.K0) GO TO 44
      IF(NP(NKA).EQ.0) GO TO 44
      A(ID+1)=SU+A(ID+1)
      44 NJ1=NP(NJA)
      IF(NJ1.EQ.0) GO TO 46
      A(ID+NJ1-NN)=F+A(ID+NJ1-NN)
      46 RHS(IR)=RES
      BB(IR)=B
      HH(IR)=H
C *****
C CALL SUBROUTINE RESI FOR MULTI-AQUIFER OPTION ONLY IF ROW,COLUMN OF
C * DO LOOP HAS A CORRESPONDING MULTI-AQUIFER NODE SET.
C * FOR NORTHERN MIDWEST STUDY TOP LAYER IS CONSTANT HEAD SO BY-PASS
C * LAYER K0. FOR GENERAL APPLICATION LAYER K0 MAY BE INCLUDED.
      IF(IOPEN.NE.ICHK(12)) GO TO 48
      IF(K.EQ.K0) GO TO 48
      IF(I.EQ.IL(IWELL).AND.J.EQ.JL(IWELL)) GO TO 4777
      GO TO 48
      4777 CALL RESI(IO,J0,K0,PHI,T,I,J,K,RESO,IWELL,IFLAG)
      RHS(IR)=RHS(IR)-RESO
      48 CONTINUE
      IF(IFLAG.NE.0) IWELL=IWELL+1
C *****
      47 CONTINUE
      IAO=NB(I-1)*MBW+1
      NEQS=NB(I)-NB(I-1)
      IF(NEQS.LT.2) GO TO 60

```

# APPENDIX F (Continued)

```

C  DECOMPOSE A MATRIX FOR SLICE                                SP31280
    CALL BAND(A(IAO),RHS,NEQS,MBW,1)                          SP31290
    60 CONTINUE                                                SP31300
CBBBBBBBBBBBBBBB END OF SLICE LOOP FOR 'A' MATRIX AND RHS TERM BBBBBBBBBBBBBB00000000
CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC00000000
    ITT=-1                                                    SP31310
C  BEGIN ITERATIONS                                           SP31320
C    SOLVE FOR HEAD DIFFERENCE AFTER PLACING B AND H COEFFICIENTS ON 00000000
C    RIGHT SIDE                                                00000000
    DO 120 IT=1,ITMX1                                         SP31330
    ITT=ITT+1                                                  SP31340
    TEST3(IT)=0.0                                              SP31350
    TEST=0.0                                                  SP31360
    BIG=0.0                                                    SP31370
    DO 110 I=2,I1                                              SP31380
    NBI1=NB(I-1)                                              SP31390
C  COMPLETE RHS TERMS                                         SP31400
    DO 90 J=2,J1                                              SP31410
    DO 90 K=1,K0                                              SP31420
    N=I+JPOS(J)+KPOS(K)                                       SP31430
    IR=NP(N)                                                  SP31440
    IF(IR.EQ.0) GO TO 90                                       SP31450
    IRT=IR+NBI1                                               SP31460
    RSMAL(IR)=RHS(IRT)-BB(IRT)*DH(N-1)-HH(IRT)*DH(N+1)      SP31470
    90 CONTINUE                                               SP31480
    IAO=NBI1*MBW+1                                           SP31490
    NEQS=NB(I)-NBI1                                           SP31500
    IF(NEQS.LT.2) GO TO 95                                     SP31510
C  SOLVE FOR DH FOR SLICE                                     SP31520
    CALL BAND(A(IAO),RSMAL,NEQS,MBW,2)                       SP31530
    GO TO 96                                                  SP31540
    95 IF(NEQS.LT.1) GO TO 110                                 SP31550
    RSMAL(1)=RSMAL(1)/A(IAO)                                  SP31560
    96 CONTINUE                                               SP31570
C  COMPLETE OVERRELAXATION                                     SP31580
    DO 100 J=2,J1                                             SP31590
    DO 100 K=1,K0                                             SP31600
    N=I+JPOS(J)+KPOS(K)                                       SP31610
    NN=NP(N)                                                  SP31620
    IF(NN.EQ.0) GO TO 100                                       SP31630
    DIFF=WO*(RSMAL(NN)-DH(N))                                  SP31640
    DH(N)=DH(N)+DIFF                                           SP31650
    TCHK=ABS(DIFF)                                             SP31660
    IF(TCHK.GT.ABS(BIG)) BIG=DIFF                              SP31670
    100 CONTINUE                                              SP31680
    110 CONTINUE                                              SP31690
C *****00000000
C * ALL SLICES FOR THIS ITERATION ARE COMPLETED. NOW TEST FOR CLOSURE. *00000000
C * BIG IS LARGEST CHANGE IN HEAD DIFFERENCES AND ERR IS CLOSURE.    *00000000
C *****00000000

```

# APPENDIX F (Continued)

C TEST TERMINATION CONDITIONS	SP31700
IF(ABS(BIG).GT.ERR) TEST=1.	SP31710
TEST3(IT)=BIG	SP31720
IF (TEST.EQ.0.) GO TO 130	SP31730
22 IF(MOD(IT,LENGTH).NE.0) GO TO 120	SP31740
C *****0000000	
C COMPUTE WATT'S CORRECTION	SP31750
C NOTE: IF THERE ARE ANY CONSTANT HEAD NODES, IN A SLICE, CORRECTION	SP31751
C IS NOT APPLIED. IN THAT TOP LAYER OF NORTHERN-MIDWEST MODEL	SP31752
C CONSISTS OF CONSTANT HEAD NODES WATT'S CORRECTION WAS NOT USED.	SP31753
ID=-1	SP31760
NEQS=0	SP31770
DO 112 I=2,I1	SP31780
NB11=NB(I-1)	SP31790
IF(NB(I).EQ.NB11.OR.ICH(I).EQ.1) GO TO 112	SP31800
SUM1=0.	SP31810
SUM2=0.	SP31820
SUM3=0.	SP31830
DO 111 J=2,J1	SP31840
DO 111 K=1,K0	SP31850
N=I+JPOS(J)+KPOS(K)	SP31860
NN=NP(N)	SP31870
IF(NN.EQ.0) GO TO 111	SP31880
IRT=NN+NB11	SP31890
NIA=N+1	SP31900
NIB=N-1	SP31910
VX=DELX(J)	SP31920
VY=DELY(I)	SP31930
VQ=VX*VY	SP31940
IF(IEQN.NE.ICHK(11)) GO TO 23	SP31950
DZ=DELZ(K)	SP31960
VX=VX*DZ	SP31970
VQ=VQ*DZ	SP31980
23 B=TC(NIB)*VX	SP31990
H=TC(N)*VX	SP32000
25 RHO=S(N)/DELT	SP32010
E=-B-H-RHO*VQ	SP32020
SUM1=SUM1+E	SP32030
RES=RHS(IRT)-B*DH(NIB)-E*DH(N)-H*DH(NIA)	SP32040
SUM3=SUM3+RES	SP32050
SUM2=SUM2+H	SP32060
111 CONTINUE	SP32070
NEQS=NEQS+1	SP32080
ID=ID+2	SP32090
AA(ID)=SUM1	SP32100
AA(ID+1)=SUM2	SP32110
IF(ICH(I+1).EQ.1) AA(ID+1)=0.	SP32120
RSMAL(NEQS)=SUM3	SP32130
112 CONTINUE	SP32140
IF(NEQS.EQ.0) GO TO 120	SP32150
IF(NEQS.LT.2) GO TO 115	SP32160

## APPENDIX F (Continued)

CALL BAND(AA,RSMAL,NEQS,2,1)	SP32170
CALL BAND(AA,RSMAL,NEQS,2,2)	SP32180
GO TO 116	SP32190
115 RSMAL(1)=RSMAL(1)/AA(1)	SP32200
116 NEQS=0	SP32210
DO 114 I=2,I1	SP32220
NB11=NB(I-1)	SP32230
IF(NB(I).EQ.NB11.OR.ICH(I).EQ.1) GO TO 114	SP32240
NEQS=NEQS+1	SP32250
DO 113 J=2,J1	SP32260
DO 113 K=1,K0	SP32270
N=I+JPOS(J)+KPOS(K)	SP32280
IF(NP(N).EQ.0) GO TO 113	SP32290
DH(N)=DH(N)+RSMAL(NEQS)	SP32300
113 CONTINUE	SP32310
114 CONTINUE	SP32320
C ***** END OF WATT'S CORRECTION SEGMENT OF CODE *****	0000000
120 CONTINUE	SP32330
CCCCCCCCCCCCCCCCCCCC END ITERATION LOOP CCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCCC	00000000
130 CONTINUE	SP32340
C FORM FINAL HEADS FOR THIS TIME STEP BY ADDING HEAD DIFFERENCES TO	0000000
C HEADS OF PREVIOUS TIME STEP.	0000000
DO 140 I=2,I1	SP32350
DO 140 J=2,J1	SP32360
DO 140 K=1,K0	SP32370
N=I+JPOS(J)+KPOS(K)	SP32380
PHI(N)=PHI(N)+DH(N)	SP32390
140 CONTINUE	SP32400
IT =ITT	SP32410
IF(TEST.EQ.0.) RETURN	SP32420
IT=ITT+1	SP32430
WRITE(6,220)	SP32440
CALL OUTPUT	SP32450
150 RETURN	SP32460
220 FORMAT ('OEXCEEDED PERMITTED NUMBER OF ITERATIONS'/' ',39('*'))	SP32470
END	SP32480
SUBROUTINE BAND (A,B,N,IHB,KKK)	BND0011
DIMENSION A(1),B(1)	BND0012
NM1=N-1	BND0013
IHB1=IHB-1	BND0014
IF(KKK.EQ.2) GO TO 30	BND0015
C DECOMPOSE A	BND0016
C	BND0017
ID=1-IHB	BND0018
DO 20 I=1,NM1	BND0019
ID=ID+IHB	BND0020
C1=1./A(ID)	BND0021
LD=ID	BND0022
L=I	BND0023
DO 15 J=1,IHB1	BND0024
L=L+1	BND0025
LD=LD+IHB	BND0026
IF(L.GT.N) GO TO 20	BND0027



# APPENDIX F (Continued)

IB=ID+J	BND0028
C=A( IB)*C1	BND0029
LB =LD-1	BND0030
DO 10 K=J,IHB1	BND0031
LB=LB+1	BND0032
10 A(LB)=A(LB)-C*A( ID+K)	BND0033
A( IB)=C	BND0034
15 CONTINUE	BND0035
20 CONTINUE	BND0036
GO TO 80	BND0037
30 CONTINUE	BND0038
C MODIFY RHS	BND0039
ID=1-IHB	BND0040
DO 50 I=1,NM1	BND0041
ID=ID+IHB	BND0042
C=B( I)	BND0043
B( I)=C/A( ID)	BND0044
IREW=MINO ( IHB1,N-I)	BND0045
DO 50 L=1,IREW	BND0046
K=I+L	BND0047
50 B(K)=B(K)-A( ID+L)*C	BND0048
ID=ID+IHB	BND0049
B(N)=B(N)/A( ID)	BND0050
C BACK SUBSTITUTE	BND0051
DO 70 I=1,NM1	BND0052
ID=ID-IHB	BND0053
L=N-I	BND0054
SUM=0.	BND0055
IREW=MINO( IHB1,N-L)	BND0056
DO 60 J=1,IREW	BND0057
60 SUM=SUM-A( ID+J)*B( L+J)	BND0058
70 B(L)=B(L)+SUM	BND0059
80 RETURN	BND0060
END	BND0061

# APPENDIX G.--SUBROUTINES MULT, RESI, AND ALOCAT

```

SUBROUTINE MULT(NWEL,A,NB,MW,JPOS,KPOS,T,NP,KO)
C *****MUL0010
C * SUBROUTINE TO COMPUTE AND LOAD MULTIAQUIFER COEFFICIENTS INTO *MUL0020
C * COMPRESSED COEFFICIENT MATRIX 'A'. COEFFICIENTS ARE THOSE TERMS OF *MUL0030
C * MULTIAQUIFER FORMULATION THAT MULTIPLY HEAD OF LAYERS OPEN TO A *MUL0040
C * WELL OR WELLS. *MUL0050
C *****MUL0060
C *****MUL0070
    PARAMETER MW=365,MJ=8,MQ=4
    DIMENSION A(1),NB(1),JPOS(1),KPOS(1),T(1),NP(1)
    COMMON/WIS/IL(MW),JL(MW),NWL(MW),KLL(MQ,MW,MJ),RWL(MW),QWL(MJ,
1MW),C2(MJ,MW),C1(MQ,MW),NU(MJ,MW),NR(MQ,MW),PI2
C *****MUL0080
C *****MUL0090
C *****MUL0100
C *****MUL0110
C *****MUL0120
    DO 100 L=1,NWEL
    I=IL(L)
    J=JL(L)
    KT=MQ
    JA=NWL(L)
    DO 700 JJ=1,JA
    DO 70 M=1,KT
    K=KLL(M,L,JJ)
    IF(K.EQ.0) GO TO 70
    N=I+JPOS(J)+KPOS(K)
    ID=(NB(I-1)+NP(N)-1)*MW+1
    A(ID)=A(ID)-PI2*C1(K,L)*NU(JJ,L)*T(N)
    DO 60 KL=M,KT
    K1=KLL(KL,L,JJ)
    IF(K1.EQ.0) GO TO 60
    N1=I+JPOS(J)+KPOS(K1)
    IB=ID+K1-K
    A(IB)=A(IB)+C1(K,L)*PI2*T(N)*NU(JJ,L)*T(N1)*C1(K1,L)*C2(JJ,L)
    60 CONTINUE
    70 CONTINUE
    700 CONTINUE
    100 CONTINUE
    RETURN
    END
SUBROUTINE RESI(IO,JO,KO,PHI,T,I,J,K,RESO,IWELL,IFLAG)
C *****RES0010
C * SUBROUTINE TO COMPUTE THE RIGHT HAND SIDE CONTRIBUTION OF MULTI- *RES0020
C * AQUIFER COEFFICIENTS (RHS TERM OF SUBROUTINE SOLVE). THE CONTRIBU- *RES0030
C * TION IS OBTAINED BY MULTIPLYING THE COEFFICIENTS BY CORRESPONDING *RES0040
C * HEADS PREVAILING AT THE END OF THE PREVIOUS TIME STEP AND SUMMING *RES0050
C * THE PRODUCTS. THE SUM TERMED RESO WHEN PASSED BACK TO SUBROUTINE *RES0060
C * SOLVE IS THEN ADDED TO THE RHS TERM FOR THE NODE. *RES0070
C *****RES0080
C *****RES0090
    REAL *8PHI
    PARAMETER MW=365,MJ=8,MQ=4
    DIMENSION PHI(IO,JO,KO),T(IO,JO,KO)
    COMMON/WIS/IL(MW),JL(MW),NWL(MW),KLL(MQ,MW,MJ),RWL(MW),QWL(MJ,
1MW),C2(MJ,MW),C1(MQ,MW),NU(MJ,MW),NR(MQ,MW),PI2
C *****RES0100
C *****RES0110
C *****RES0120
C *****RES0130
C *****RES0140
C *****RES0150

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# APPENDIX G (Continued)

RES2=0.0	RES0160
RES1=0.0	RES0170
KK=MQ	RES0190
KW=0	RES0200
L=IWELL	RES0210
JA=NWL(L)	RES0220
C WELL IS A COUNTER FOR PUMPED NODE(MULTIAQUIFER)	RES0230
C KW COUNTS THE NUMBER OF TIMES LAYER K IS TAPPED	RES0240
C IF LAYER K IS NOT TAPPED LOOP TO 70	RES0250
DO 70 JJ=1,JA	RES0260
SUMZ=0.0	RES0270
DO 60 KL=1,KK	RES0280
K1=KLL(KL,L,JJ)	RES0290
IF(K1.EQ.K) KW=KW+1	RES0300
IF(K1.EQ.0.AND.KL.EQ.K) GO TO 70	RES0310
IF(K1.EQ.0) GO TO 60	RES0320
SUMZ=SUMZ+PHI(I,J,K1)*C1(K1,L)*T(I,J,K1)	RES0330
60 CONTINUE	RES0340
RES2=RES2+PI2*C1(K,L)*NU(JJ,L)*C2(JJ,L)*T(I,J,K)*SUMZ	RES0350
70 CONTINUE	RES0360
IF(KW.EQ.0) GO TO 75	RES0370
RES1=-C1(K,L)*PI2*T(I,J,K)*PHI(I,J,K)*NR(K,L)	RES0380
75 RESO=RES2+RES1	RES0390
IFLAG=1	RES0400
RETURN	RES0410
END	RES0420
SUBROUTINE ALOCAT(PHI,OLD,T,S,TK,DELX,DELY,DELZ,IO,JO,KO,K5,HW,Q, ALO0010	
1NS1,II,I,J,TC,TR,QRE) ALO0020	
C *****AL00030	
C * SUBROUTINE TO COMPUTE AND PRINT FLOW COMPONENTS FOR LAYERS OF NODE *AL00040	
C * SETS CONTAINING MULTIAQUIFER WELLS AND TO COMPUTE BALANCE OF FLOWS *AL00050	
C * SUBROUTINE PRIMARILY USED TO TEST AND DEBUG MULTIAQUIFER CHANGES *AL00060	
C * IN SUBROUTINE SOLVE AND CODE OF SUBROUTINES MULT AND RES1. *AL00070	
C * SUBROUTINE ALOCAT IS CALLED FROM SUBROUTINE CHECKI(STATEMENT NUMBER*AL00080	
C * CHK1578) AND IS PRESENTLY INVOKED FOR THE LAST STRESS PERIOD ONLY. *AL00090	
C * QL IS SUM OF LATERAL FLOW BETWEEN NODE AND ADJACENT NODES. *AL00100	
C * QS IS FLOW IN/OUT OF STORAGE *AL00110	
C * QQ IS SUM OF FLOW BETWEEN LAYER K AND LAYERS K-1,AND K+1 *AL00120	
C * QOUT IS FLOW BETWEEN NODE AND WELL *AL00130	
C * QQQ IS QOUT IN CUBIC FEET PER DAY *AL00140	
C * PST IS VARIABLE DENSITY SOURCE TERM FOR NODE *AL00150	
C * DIFF IS SUM OF ALL FLOWS I.E. IT IS NODAL BALANCE *AL00160	
C *****AL00170	
REAL *8PHI,P ALO0180	
PARAMETER MW=365,MJ=8,MQ=4 ALO0190	
DIMENSION PHI(IO,JO,KO),OLD(IO,JO,KO),T(IO,JO,KO),S(IO,JO,KO), ALO0200	
1TK(IO,JO,K5),DELX(JO),DELY(IO),DELZ(KO),TC(IO,JO,KO),TR(IO,JO,KO) ALO0210	
DIMENSION QV(6),HW(MJ),QRE(IO,JO,KO) ALO0220	
COMMON/WIS/IL(MW),JL(MW),NWL(MW),KLL(MQ,MW,MJ),RWL(MW),QWL(MJ, ALO0230	
1MW),C2(MJ,MW),C1(MQ,MW),NU(MJ,MW),NR(MQ,MW),PI2 ALO0240	
COMMON /SPARAM/ TMAX,CDLT,DELT,ERR,TEST,SUM,SUMP,QR ALO0250	
C ALO0260	

# APPENDIX G (Continued)

AREA=DELX(J)*DELY(I)	AL00270
KK=MQ	AL00290
WRITE(6,21)	AL00300
21 FORMAT(1H0,'        NODE')	AL00310
K6=K0+1	AL00320
DO 70 L=1,K6	AL00330
70 QV(L)=0.0	AL00340
DO 71 L1=2,K0	AL00350
L=L1-1	AL00360
QV(L1)=TK(I,J,L)*(PHI(I,J,L)-PHI(I,J,L1))	AL00370
71 QV(L1)=+QV(L1)*AREA	AL00380
DO 320 L=1,KK	AL00390
P=PHI(I,J,L)	AL00400
QA=TR(I,J-1,L)*(PHI(I,J-1,L)-P)*DELY(I)	AL00410
QB=TR(I,J,L)*(PHI(I,J+1,L)-P)*DELY(I)	AL00420
QC=TC(I-1,J,L)*(PHI(I-1,J,L)-P)*DELX(J)	AL00430
QD=TC(I,J,L)*(PHI(I+1,J,L)-P)*DELX(J)	AL00440
QL=(QA+QB+QC+QD)	AL00450
QS=S(I,J,L)*(P-OLD(I,J,L))	AL00460
QS=-QS*AREA/DELT	AL00470
IF(S(I,J,L).LT.0.0) QS=999999.	AL00480
QOUT=0.0	AL00490
JA=NWL(II)	AL00500
DO 79 JJ=1,JA	AL00510
K=KLL(L,II,JJ)	AL00520
IF(K.EQ.0) GO TO 79	AL00530
QOUT=QOUT+PI2*C1(K,II)*NU(JJ,II)*T(I,J,K)*(P-HW(JJ))	AL00540
79 CONTINUE	AL00550
QQQ=QOUT*86400.	AL00560
QOUT=-QOUT	AL00570
QQ=QV(L)-QV(L+1)	AL00580
PST=QRE(I,J,L)	AL00590
DIFF=QL+QS+QQ+QOUT+PST	AL00600
IF(S(I,J,L).LT.0.0) DIFF=999.99999	AL00610
699 CONTINUE	AL00620
WRITE(6,20)I,J,L,QL,QS,QQ,QOUT,PST,QQQ,DIFF	AL00630
20 FORMAT(2X,I2,',',I2,',',I2,6(3X,E14.8),6X,F12.5)	AL00640
320 CONTINUE	AL00650
END	AL00660