

PRECIPITATION

PRECIPITATION, STREAMFLOW, AND BASE FLOW  
IN WEST-CENTRAL TEXAS, DECEMBER 1974 THROUGH MARCH 1977

INTRODUCTION

Precipitation, streamflow, and base-flow data were analyzed for December 1974 through March 1977 as a part of the Edwards-Trinity Regional Aquifer-Systems Analysis (RASA) project (Bush, 1986). These data are key components in water-budget calculations for estimating recharge in the area underlain by the Edwards-Trinity aquifer system in west-central Texas (fig. 1). The period of record analyzed corresponds to the calibration period of a digital ground-water flow model currently (1988) being constructed. In a simple system the appropriate equations for several components of the water budget are:

$R = P - ET - SRO,$   
 $R = BF + DR,$   
 $TS = BF + SRO$

(1)  
(2)  
(3)

where

- R is shallow recharge, in inches;
- P is precipitation, in inches;
- ET is evapotranspiration, in inches;
- SRO is surface runoff, in inches;
- BF is base flow, in inches;
- DR is deep recharge, in inches; and
- TS is total streamflow, in inches.

Precipitation and the precipitation departure from normal (1895-1985) (Karl and Knight, 1985a, b, c, d) are presented on sheet 1. Information about total streamflow and base flow is presented on sheet 2.

PHYSIOGRAPHY AND HYDROLOGIC SETTING

The Edwards-Trinity aquifer system in west-central Texas is bounded on the south and east by the Balcones fault zone along the southern and eastern edge of the Hill Country physiographic region (fig. 1). The Hill Country region is characterized by rolling terrain ranging in altitude from 400 to 2,400 ft (Ashworth, 1983, p. 2). The center of the study area is in the Edwards Plateau physiographic region, characterized by rolling plains to flat tableland and rugged, steep-walled canyons and draws ranging in altitude from 1,000 to 3,300 ft (Walker, 1979, p. 7). To the west, the Edwards-Trinity aquifer system is bounded by a series of mountain ranges at the center of the Trans-Pecos physiographic region. The Trans-Pecos region is characterized by the flat alluvial valley of the Pecos River in the north and east, the extensively dissected flat plateaus and mesas in the south, and the rugged, high-relief terrain of the Davis Mountains on the west. Altitudes in the Trans-Pecos region range from 1,200 ft in the south to 8,000 ft in the Davis Mountains (Rees, Rhye, and Buckner, 1980, p. 2).

The northwestern part of the study area is in the southern part of the High Plains physiographic region. The High Plains region is flat with few streams. Altitudes in the southern High Plains range from 2,600 ft in the east to 3,300 ft in the west. The northeastern part of the study area is in the Central Minerals physiographic region. The Central Minerals region is hilly with high relief and deeply incised streams. The Central Minerals region is lower in altitude than the Edwards Plateau with altitudes ranging from 1,000 ft near the Colorado and Llano Rivers to 1,800 ft near the Edwards Plateau.

The climate varies from subhumid in the east to arid in the west. There are two rainy seasons, one in spring and one in fall, in the eastern part of the study area. Intense storms may occur in spring, summer, and fall (March through November). In winter (December through February) storms are less intense. Storms in the eastern part of the area usually are widespread. In the western part of the study area, precipitation usually occurs in the summer. These summer storms may be intense, but are local in extent. Precipitation has the greatest spatial variability and the least frequency in the west. Mean annual precipitation (1951-80) throughout the study area varies from 32 inches in the east to 10 inches in the west (fig. 2).

Potential evapotranspiration is a theoretical value representing the maximum quantity of water that could be used by plants if precipitation were sufficient to supply this quantity of water to the soil. Potential evapotranspiration commonly is a mathematical function of temperature and number of daylight hours. In the study area potential evapotranspiration ranges from 36 to 46 in/yr (Geraghty and others, 1973, pl. 13). Actual evapotranspiration is much less than the theoretical value for potential evapotranspiration in the study area because precipitation is much less than potential evapotranspiration. Potential evapotranspiration can be as much as 18 to 38 in/yr greater than precipitation from east to west, respectively.

Four major river basins drain the study area: the Colorado, the Guadalupe, the Nueces, and the Rio Grande. Mean annual streamflow (1951-80) in these basins ranges from 5 in/yr on the Balcones fault zone in the east to less than 0.1 in/yr in the Rio Grande basin in the west (Walter Lear, U.S. Geological Survey, written commun., 1986).

The major aquifers of the Edwards-Trinity system occur in rocks of Cretaceous age. In the Balcones fault zone, the principal aquifer is the Edwards. In the Hill Country, the principal aquifer is the Trinity. In the Edwards Plateau and the Trans-Pecos, the principal aquifer is the Edwards-Trinity. Lithologically, these aquifers are predominantly carbonate with some basal sand formations. In the Trans-Pecos region, the Edwards-Trinity is overlain by and hydraulically connected to an alluvial aquifer near the Pecos River that consists of sediment of Cenozoic age. The High Plains aquifer to the northwest of the Edwards-Trinity consists of sand of Cenozoic age that is adjacent to and hydraulically connected to the sand of the Edwards-Trinity in the Edwards Plateau. In the Central Minerals region, the Hickory Sandstone and Ellenburger-San Saba (Mestones) aquifers consist of older rocks ranging in age from Cambrian to Ordovician. Precambrian rocks of eroded, faulted, and fractured granite also crop out in the region (Walker, 1979, table 2).

Throughout the Central Minerals, Edwards Plateau, Trans-Pecos, and northwestern part of the Hill Country physiographic regions, the Edwards-Trinity aquifer is unconfined to semiconfined. The water table varies from near land surface adjacent to some of the streams to as much as 800 ft below land surface in the mountains. In general, the water table is 100 to 200 ft below the land surface throughout most of the Edwards Plateau. Gradients of the water table indicate ground-water movement toward the headwaters of streams in the Hill Country and toward all streams in the Central Minerals, Edwards Plateau, and Trans-Pecos physiographic regions (Walker, 1979, fig. 15; Rees, Rhye, and Buckner, 1980, fig. 7). In the Edwards aquifer along the Balcones fault zone, water-table conditions exist in a narrow strip of outcrop adjacent to the Hill Country and confined conditions exist downdip from the outcrop. Gradients indicate flow from southwest to northeast along the strike of the fault zone (Macley and Smell, 1986, fig. 23).

PRECIPITATION

The Edwards-Trinity aquifer system in west-central Texas underlies all or part of several climatic divisions. The major divisions are the High Plains, the Trans-Pecos, the Edwards Plateau, and the South Central (fig. 3). Total precipitation was computed using National Weather Service rain-gage data. Thiessen polygons were constructed using monthly data; the data were summed for the 28-month calibration period and divided by 2.333 to convert the values to inches per year. Precipitation ranged from 30 to 45 in/yr in the South Central division within the study area, 13 to 41 in/yr in the Edwards Plateau division, 6 to 21 in/yr in the Trans-Pecos division within the study area, and 16 to 22 in/yr at the two stations in the High Plains division within the study area. Areas of equal 28-month mean precipitation are shown in figure 3.

Precipitation departure from normal (1951-80) was computed for the calibration period by subtracting the mean annual precipitation (fig. 2) for the 30-year period from the areas of equal 28-month mean precipitation, in inches per year (fig. 3). Precipitation ranged from 6 in/yr less than normal to 15 in/yr greater than normal in the major divisions (fig. 4). Because the calibration period is 28 months starting in December, the 28-month mean is biased by an extra winter season.

Seasonal and monthly precipitation departure from normal (1895-1985) has been determined for 344 climatic divisions for the contiguous United States by Karl and Knight (1985a, b, c, d). They analyzed precipitation data for each division and fit the data to a three-parameter Pearson Type III distribution, also known as the gamma distribution.

Seasonal departures from normal for all of Texas are shown in figure 5 for December 1974 through May 1977. Class limits, or ranges, for each category of precipitation are listed in table 1. The probability of occurrence, in percent, shown for each category in figure 5 indicates that seasonal precipitation in either the extremely wet or the extremely dry category has a less than 5-percent chance of occurrence; precipitation in either the very wet or the very dry category has a 10-percent chance of occurrence; precipitation in either the moderately wet or the moderately dry category has a 15-percent chance of occurrence; and near-normal precipitation has a 40-percent chance of occurrence. In the study area during the calibration period, precipitation ranged from near normal to extremely dry from September 1975 through February 1976. Precipitation ranged from near normal to extremely wet from March 1976 through May 1976 and from September 1976 through November 1976 (fig. 5). Winter is the driest season for the climatic divisions in the study area. During winter 1976-77, the extra season in the 28-month mean, conditions were near normal over most of the study area.

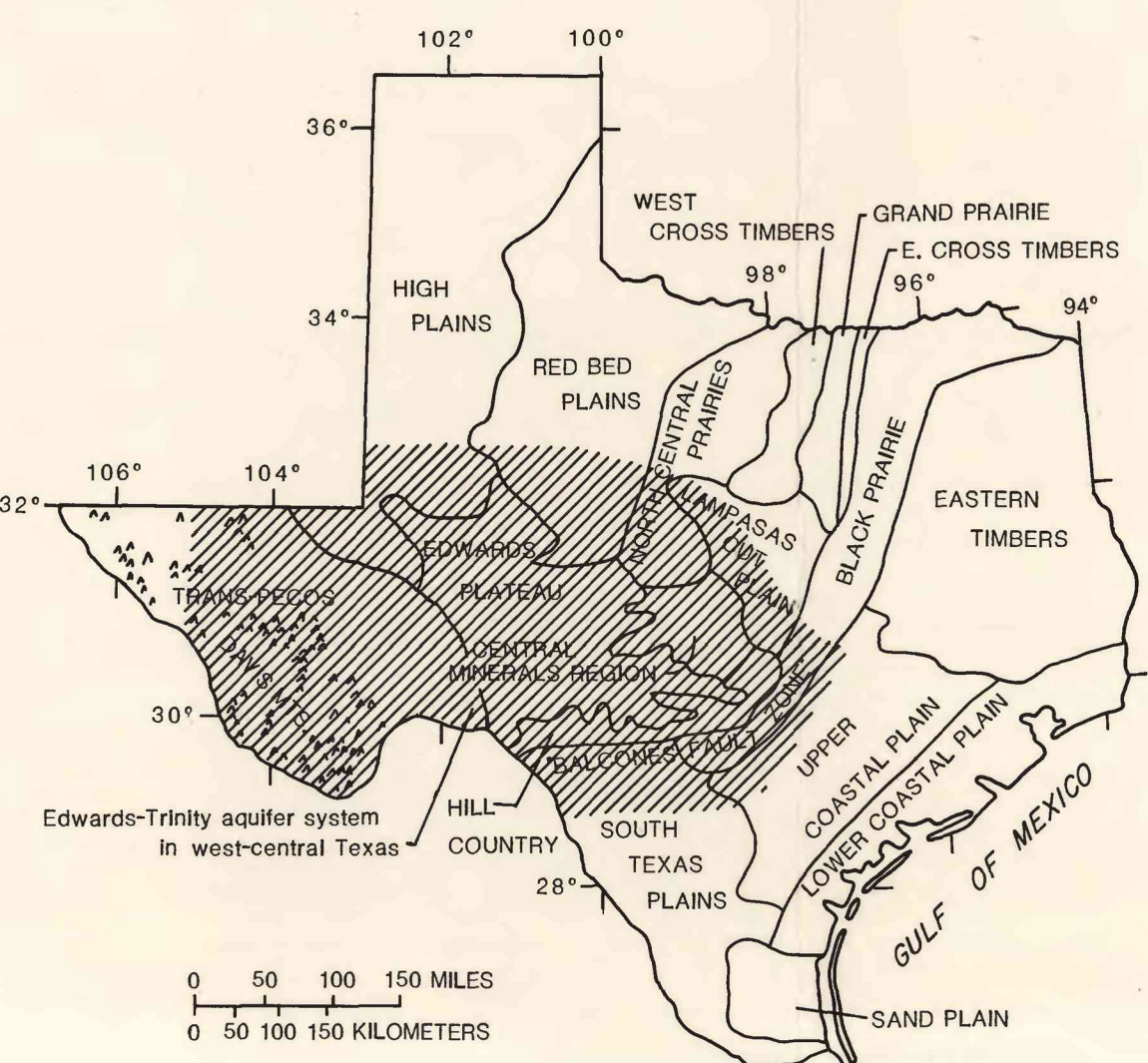


Figure 1.--Physiographic regions of Texas and location of the Edwards-Trinity aquifer system in west-central Texas. (Modified from Kier and others, 1977)

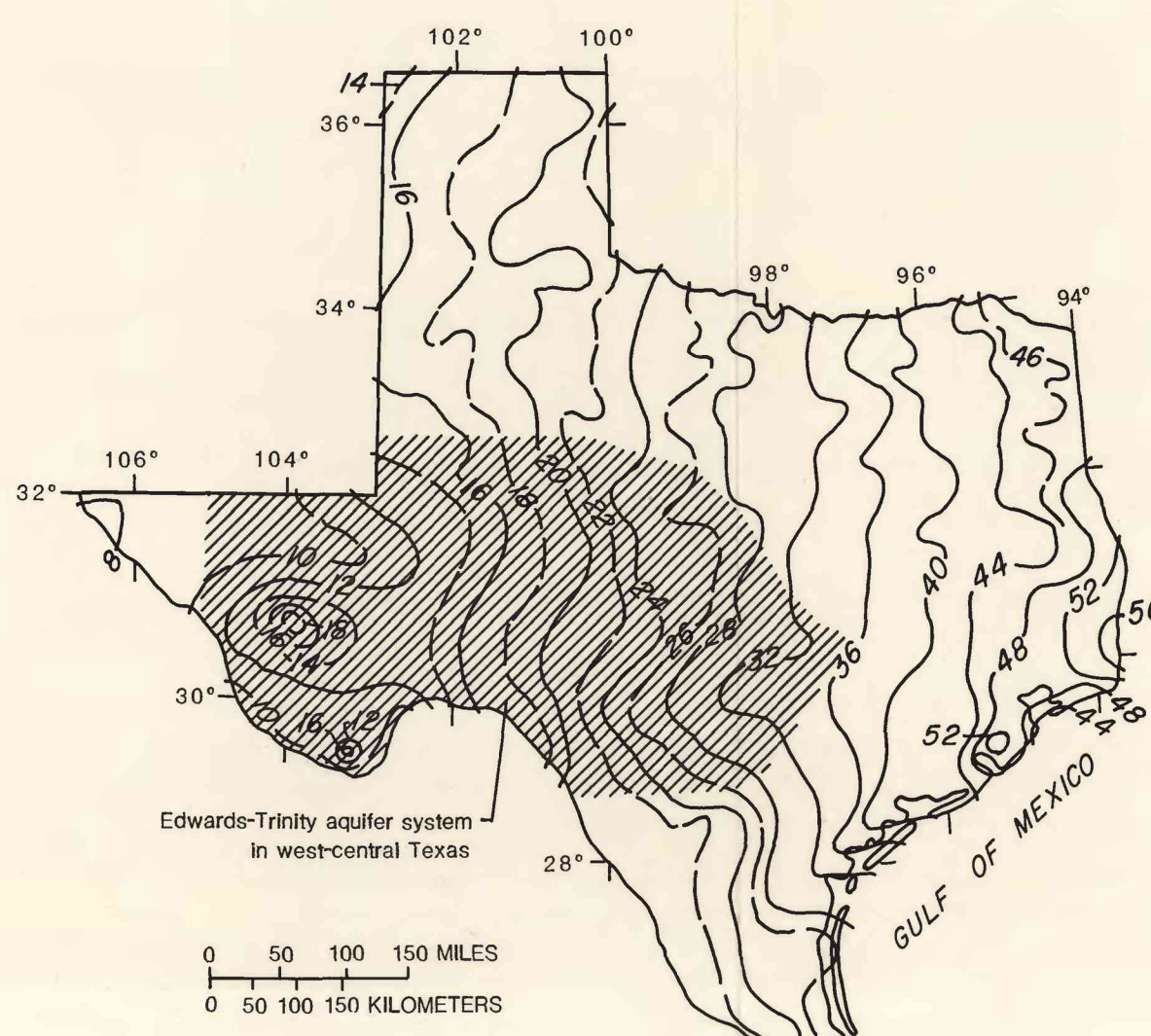


Figure 2.--Mean annual precipitation in inches, 1951-80. (From Riglio and others, 1987, fig. 11)

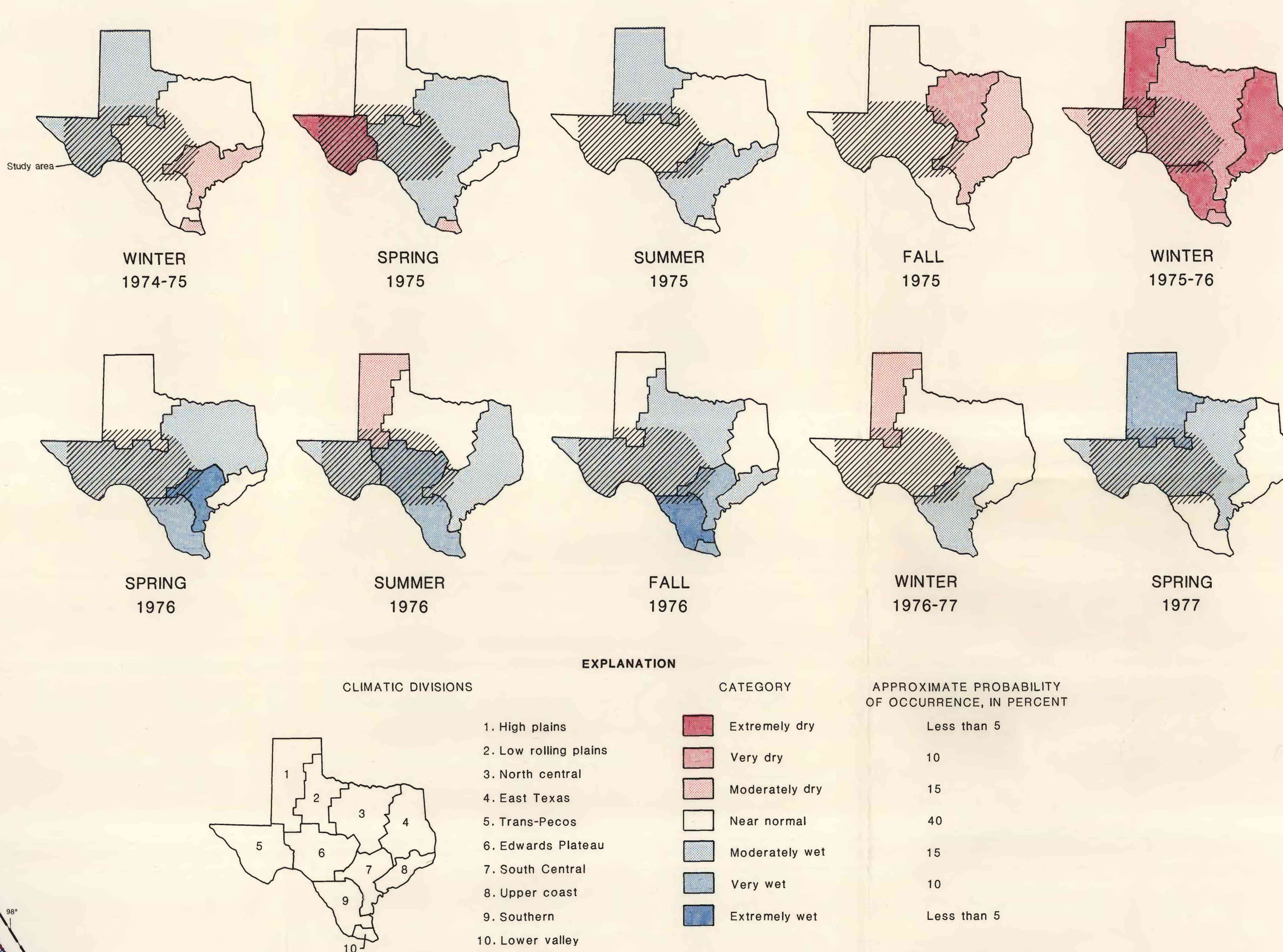


Figure 5.--Seasonal precipitation departures from normal (1895-1985) for the climatic divisions of Texas, December 1974 through May 1977 (Karl and Knight, 1985 a,b,c,d).

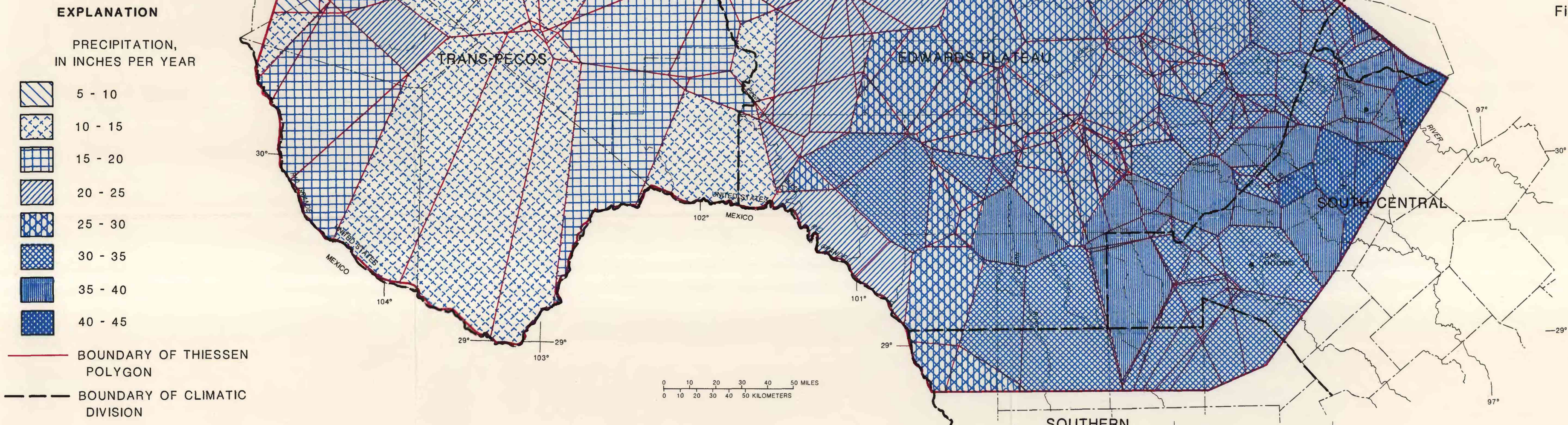


Figure 3.--Areas of equal 28-month mean precipitation in west-central Texas, December 1974 through March 1977.

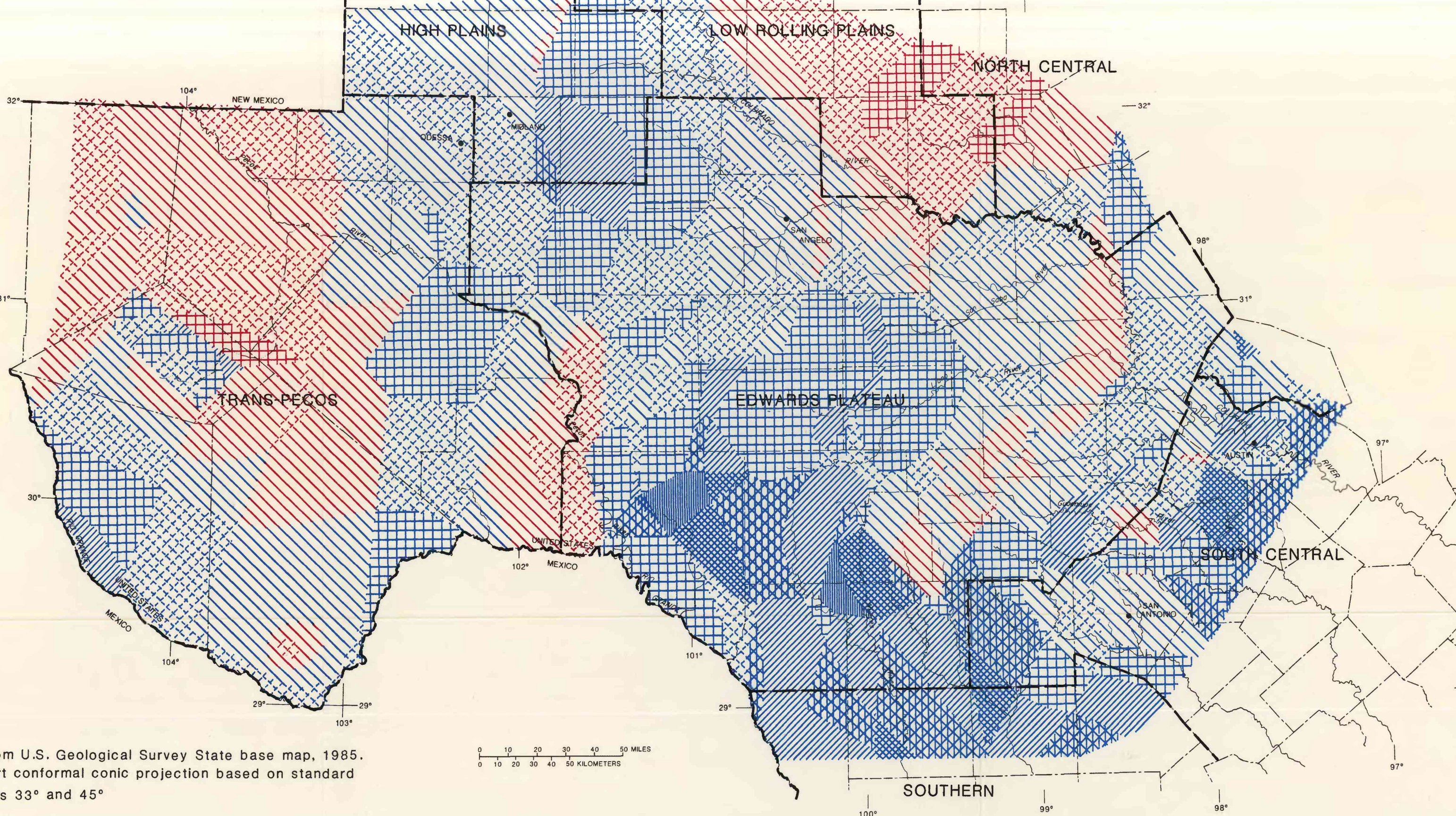


Figure 4.--Areas of equal 28-month mean precipitation departure from normal (1951-80) in west-central Texas, December 1974 through March 1977.

PRECIPITATION, STREAMFLOW, AND BASE FLOW IN WEST-CENTRAL TEXAS, DECEMBER 1974 THROUGH MARCH 1977

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## STREAMFLOW

Total streamflow equals precipitation that runs off the land surface plus precipitation that infiltrates into the ground and eventually seeps into a stream. Streamflow is gaged at streamflow-gaging stations located throughout the study area (fig. 6). Diversions and effects of regulations by dams are not represented in these gaged discharges. Footnotes for the streamflow-gaging stations described in table 2, indicate that few streams in Texas are unregulated. The total-streamflow column (table 2) includes the volume of known diversions from the stream upstream from the gage. The basins shaded in dark blue (fig. 6) are those upstream from the most upstream gage on each stream. The streamflow shown in these basins matches the computed annual flow given in table 2. The basins shaded in light blue (fig. 6) represent intervening basins between gages on streams. The numbers shown in these basins represent the gain (+) or loss (-) in streamflow between the gages. The accuracy of the numbers in Figure 6 and table 2 is variable, depending on the volume of unknown diversions. The uncolored basins in figure 6 represent basins for which evaporation from large reservoirs or large unknown diversions or both preclude meaningful estimates of total streamflow.

In general, streamflow increases in the downstream direction; therefore, the difference in streamflow at successive downstream gages represents a gain in streamflow throughout most of the study area. Where streams cross faulted, fractured, and/or karstic terrain and the water table is below the streambed, such as along the Balcones fault zone, water from the stream flows directly into the ground (see footnotes for total-streamflow column in table 2). This results in streamflow losses between successive downstream gages (the negative values shown in fig. 6) along several streams in the Balcones fault zone.

Precipitation is the major factor in the variation of streamflow from year to year. Areas of equal precipitation departure from normal (1951-80) are shown in figure 4. In general, precipitation was near normal in the northern and eastern parts of the study area and greater than normal in the southeastern part of the study area. Total streamflow varied from about 13 in/yr in the Balcones fault zone to less than 1 in/yr in the western part of the study area (table 2, fig. 6). In the Balcones fault zone, in the southeastern part of the study area, the streamflow was as much as 8 in/yr greater than normal (1951-80).

## BASE FLOW

Base flow is that part of precipitation recharged at topographically high areas that is discharged to streams at topographically low areas. Base flow also can be defined as the part of total streamflow that is not direct surface runoff. In general, direct surface runoff occurs quickly after a storm, whereas base flow is slowly released to a stream. It is this time delay in the appearance of base flow in a stream that allows the base-flow contribution to streamflow to be graphically separated from the surface-runoff contribution on a discharge hydrograph. Hydrographs were separated in this manner for selected streamflow-gaging stations; an example of this separation of a daily discharge hydrograph is shown in figure 7. Monthly mean base flow was determined and averaged for the 28-month calibration period. These base flows are tabulated in table 3.

Generally, streams with sustained base flow will have larger ratios of base flow to total streamflow. These ratios for the stations analyzed, expressed in percent in table 3, ranged from 24.6 to 89.5 percent. The largest base-flow values, from about 4 to 6 in/yr, are for basins in the southeastern part of the study area where precipitation is greatest. In the Rio Grande basin and in the upper Colorado River basin, base-flow values are not shown because they are less than 0.1 in/yr (fig. 6). As in figure 6, a distinction in values is made in figure 8 between basins in which the labeled number is the base flow for the entire contributing area upstream from the gage (dark blue basins) and basins in which the labeled number is the base flow for the intervening basin between successive gages (light blue basins).

In other parts of the United States, base flow has been related to flow-duration statistics. Flow-duration statistics are determined by ranking gaged daily discharges by magnitude and computing the percentage of time a discharge of a specific magnitude is equaled or exceeded. A flow-duration curve is the resulting cumulative frequency distribution of the daily discharges. Examples of flow-duration curves at two of the stations where base flow was determined are shown in figure 9.

In the Delmarva Peninsula of Maryland and Delaware, Cushing and others (1973) determined that mean annual base flow was approximately equal to the discharge equaled 50 percent of the time (Q50). In the outcrop area of the sand aquifers of the southeast coastal plain in Mississippi, Alabama, Georgia, and South Carolina, Stricker (1983) determined that mean annual base flow was approximately equal to the discharge equaled 60 to 65 percent of the time. However, no such relation could be established for the streams in west-central Texas because there was no correlation between any one flow-duration statistic and mean annual base flow (plots of mean base flow versus any one flow-duration statistic had considerable scatter rather than plotting in a band or a line).

The flow-duration index, defined as the square root of the ratio of the discharge equaled or exceeded 25 percent of the time to the discharge equaled or exceeded 75 percent of the time, is shown in column 7 of table 3. This index is used as an indication of the shape of the flow-duration curve. The greater the index, the steeper the curve, and the lesser the index, the flatter the curve (fig. 9). A steep flow-duration curve indicates a stream in which discharge decreases rapidly after a storm. A flat flow-duration curve indicates a stream with well-sustained base flow. Generally, the eight stations that have an index greater than 2 are on streams in which base flow is a smaller percentage of gaged flow. The average base flow for these eight stations is 45 percent of the gaged flow. The 29 stations that have an index less than 2 are on streams that have an average base flow of 65 percent of the gaged flow.

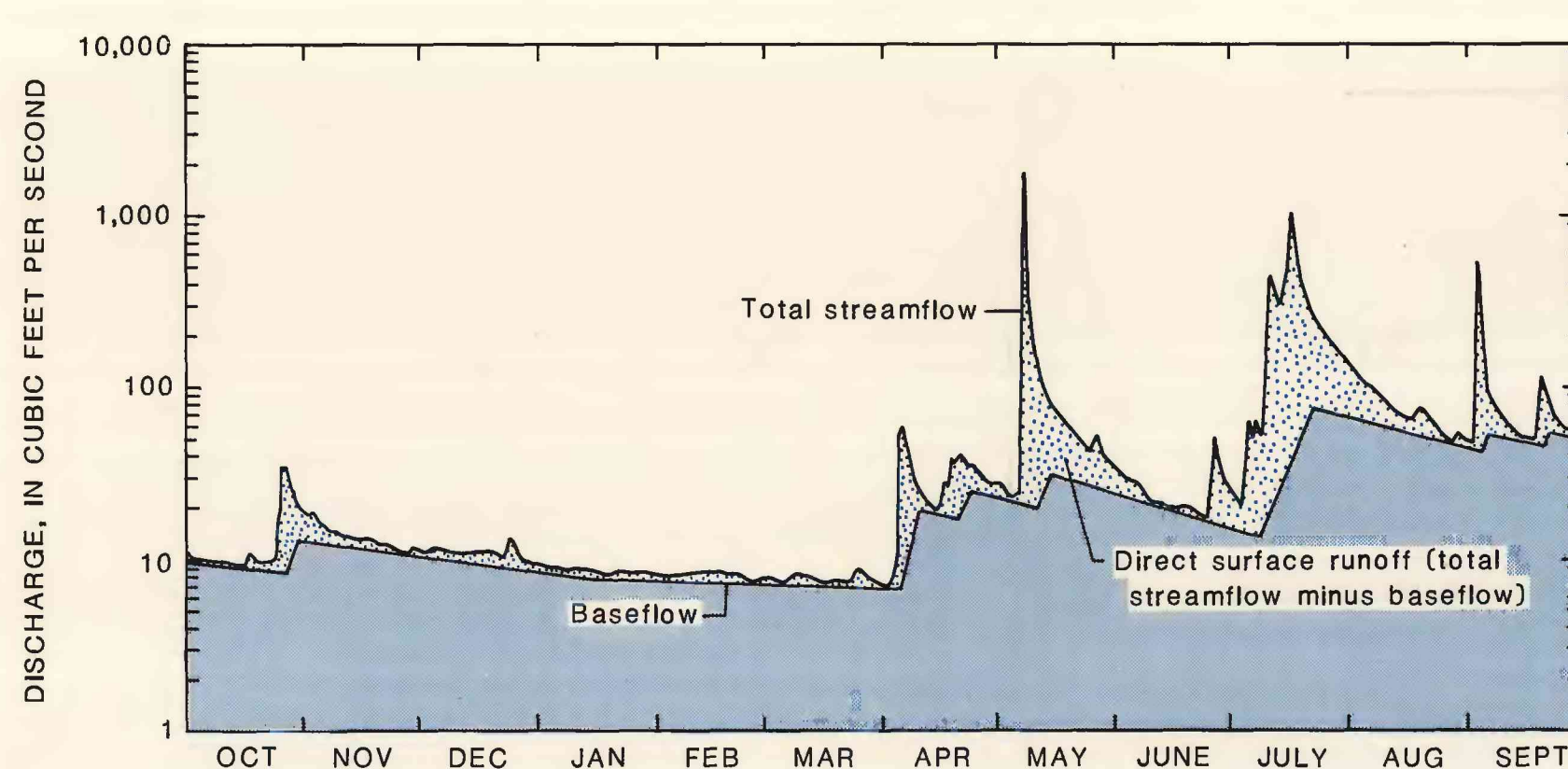


Figure 7.—Daily discharge hydrograph showing base-flow separation, station 66 (08196000), Dry Rio River near Reagan Wells, water year 1976.

Table 2.—Streamflow-gaging stations in west-central Texas used to determine base flow, December 1974 through March 1977

(in<sup>3</sup>, square miles; ft<sup>3</sup>/s, cubic feet per second; (Q25/Q75), the square root of flow with 25-percent duration divided by flow with 75-percent duration; in, inches; in/yr, inches per year. Note: Refer to table 1 for remarks about the streamflow-gaging stations.)

Refer- ence number	Station number	Station name	Contrib- uting drainage area (mi <sup>2</sup> )	Mean base flow (in/yr)	Percent of total flow	Duration index (Q25/Q75)	Base flow (in/yr)	28-month mean base flow (in/yr)
8	08107000	Elm Creek at Ballinger	450.0	12.4	28.4	3.485	0.870	0.370
9	08128000	South Concho River at Christoval	344.0	88.7	89.5	1.385	5.129	2.185
10	08128400	Middle Concho River above Tamerley	1,644.0	14.0	70.4	1.655	.269	.115
11	08130500	Spring Creek above Tamerley	405.0	17.2	65.7	1.337	1.345	.876
12	08130500	Dove Creek at Knickerbocker	218.0	34.7	83.9	1.170	5.043	2.161
13	08131400	Pecos River near San Angelo	81.1	2.5	67.2	2.974	.970	.416
14	08135000	Concho River at Ballinger	1,020.2	37.1	44.3	1.488	1.639	.687
15	08135000	Concho River near Stacy	3,234.0	116.1	49.7	1.644	1.135	.487
16	08135000	Concho River at Winnell	1,231.0	102.4	98.3	1.897	.766	.326
26	08145000	San Saba River at Menard	1,128.0	64.4	74.8	1.281	1.808	.774
27	08145000	Brady Creek near Eden	101.0	.8	53.3	1.526	.280	.107
28	08145000	San Saba River at San Saba	3,038.0	165.7	73.9	1.408	1.726	.740
29	08145000	San Saba River near Mason	3,378.0	84.4	1.803	1.305	.818	
31	08145000	North Llano River near Junction	814.0	88.8	69.4	1.388	2.382	1.021
32	08150000	Llano River at Llano	1,046.0	162.2	81.5	1.182	3.244	1.322
33	08150000	Llano River near Mason	3,342.0	239.8	72.3	1.347	2.174	.932
34	08150000	Beaver Creek near Mason	218.0	4.4	24.6	1.646	.841	.275
35	08150000	Llano River at Llano	4,192.0	32.5	39.6	2.088	2.973	1.274
36	08150000	Sandy Creek near Kingsland	346.0	103.4	56.1	1.725	3.586	2.308
37	08150000	Pedernales River near Johnson City	901.0	15.4	56.1	2.408	4.438	2.768
39	08151500	Walnut Creek at Webberville Rd., Austin	51.3	10.4	4.4	7.271	4.233	1.886
40	08151500	Wilbarger Creek near Pflugerville	169.0	27.7	82.5	1.276	5.318	2.237
41	08155000	North Fork Guadalupe River near Hunt	169.0	16.3	76.0	1.720	1.720	.888
42	08155000	Johnson Creek near Ingram	114.0	24.3	79.6	1.264	6.740	2.888
44	08170000	Guadalupe River at Comfort	839.0	178.4	65.1	1.419	6.735	2.886
45	08170000	Guadalupe River near Spring Branch	1,315.0	300.3	61.8	1.601	7.952	3.458
46	08170000	Blanco River at Wimberly	385.0	140.1	86.0	1.487	5.352	2.178
47	08170000	Salado Creek (lower station) at San Antonio	189.0	26.9	48.9	1.420	4.508	1.932
48	08170000	Medina River near Pipe Creek	474.0	157.8	68.4	1.648	10.532	4.514
49	08171000	Red Bluff Creek near Pipe Creek	56.3	6.5	31.2	51.001	3.693	1.866
50	08171000	Medina River at Laguna	764.0	14.8	79.2	1.507	6.087	2.600
51	08171000	Frio River at Canyon	405.0	107.9	72.4	1.352	8.438	3.612
52	08171000	Dry Frio River near Reagan Wells	117.0	27.9	69.7	1.820	7.533	3.238
53	08171000	Salado River near Tarpier	236.0	86.1	82.7	1.718	8.612	3.481
70	08200000	Hondo Creek near Tarpier	86.2	37.9	55.8	2.058	13.408	5.961
72	08201500	Soto Creek at Miller Ranch near Utopia	43.1	13.6	61.0	2.059	6.848	4.283

a/ Flow at station partly regulated by a reservoir thought to be far enough upstream not to interfere with separation of direct surface runoff from base flow.

## STREAMFLOW

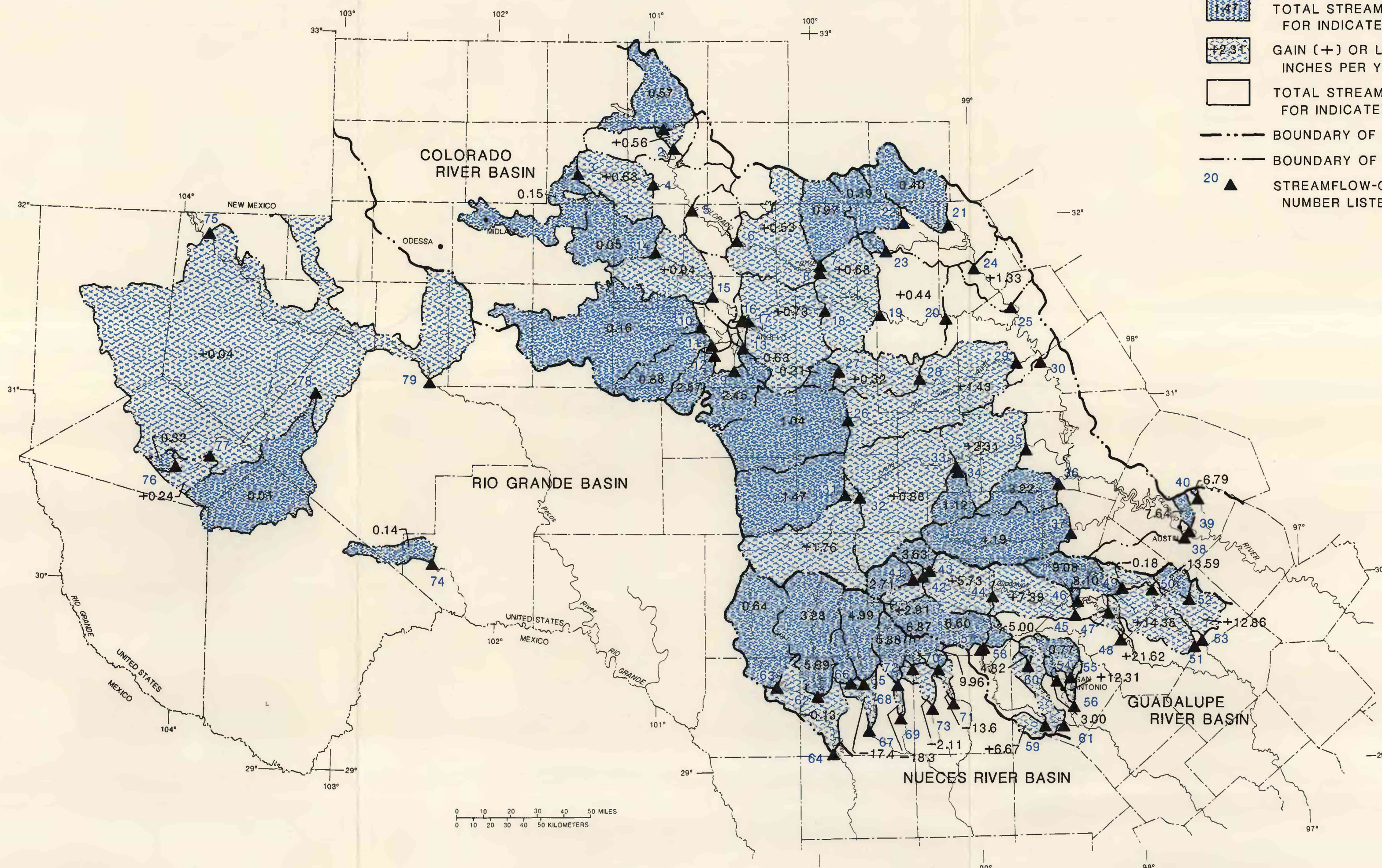


Figure 6.—Total 28-month mean streamflow values by drainage basin, December 1974 through March 1977.

## BASE FLOW

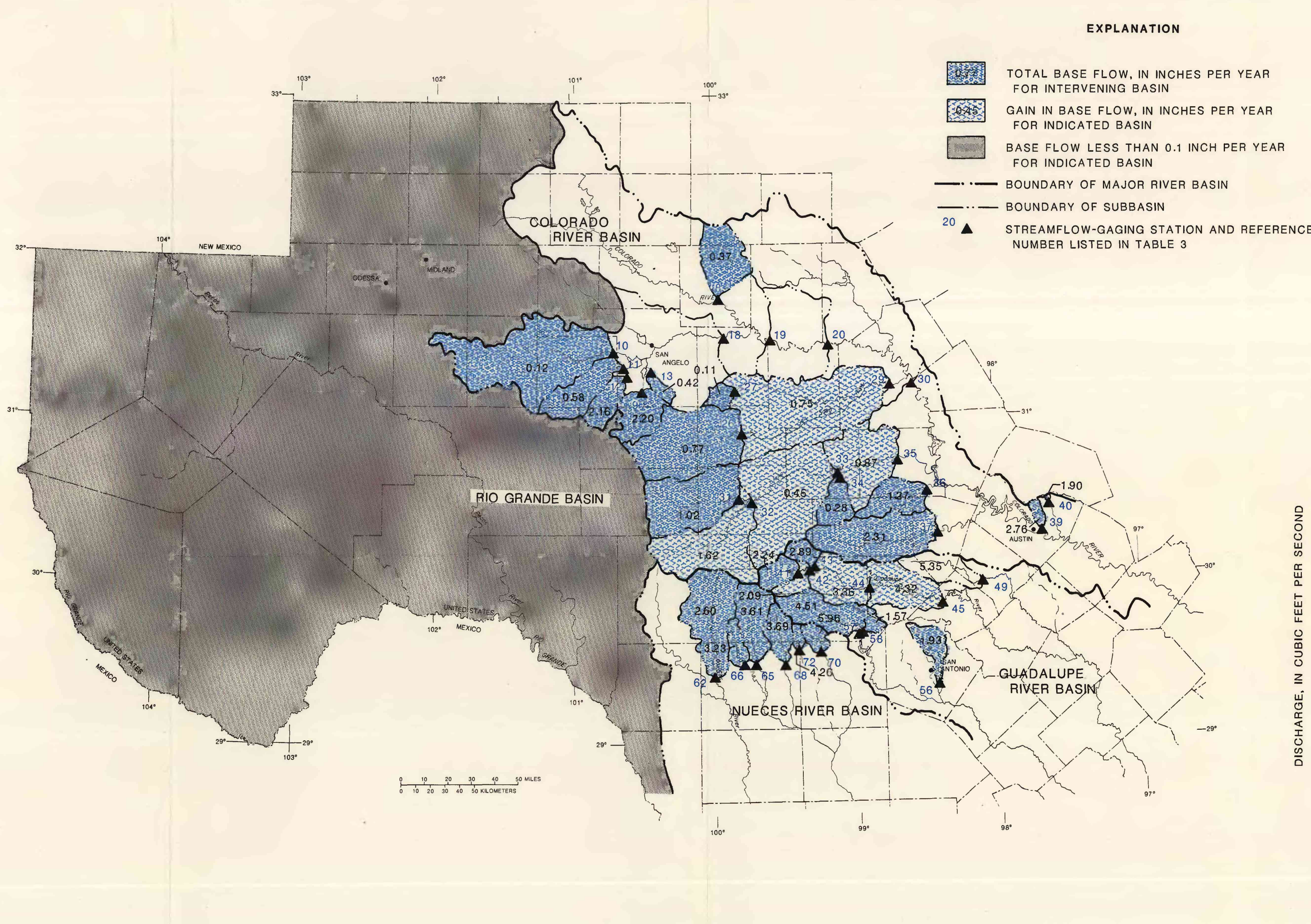


Figure 8.—Twenty eight-month mean base-flow values by drainage basin, December 1977 through March 1977.

## EXPLANATION

- TOTAL STREAMFLOW, IN INCHES PER YEAR, FOR INDICATED BASIN
- GAIN (+) OR LOSS (-) IN STREAMFLOW, IN INCHES PER YEAR FOR INTERVENING BASIN
- BOUNDARY OF MAJOR RIVER BASIN
- BOUNDARY OF SUBBASIN
- STREAMFLOW-GAGING STATION AND REFERENCE NUMBER LISTED IN TABLE 2

Table 2.—Streamflow-gaging stations in west-central Texas used to determine total runoff, December 1974 through March 1977

(in<sup>3</sup>, square miles; acre-ft, acre-feet; ft<sup>3</sup>/s, cubic feet per second; in, inches; in/yr, inches per year. Total volume-amounts are from stations not regulated by a lake or a reservoir larger than 25,000 acre-ft, but flow may be affected by small reservoirs, unless otherwise indicated.)

Refer-	Station	Station name	Contrib-	Total	Mean	Total	28-Month
ence	number		uting	volume	flow	runoff	mean
number			drainage	(acre-ft)	(ft <sup>3</sup> /s)	(in)	runoff
			area				
			(mi <sup>2</sup> )				(in/yr)
1	08120700	Colorado River near Cuthbert	a/ 549.0	b/ 38,894.0	23.0	1.328	0.568
2	08120700	Colorado River at Colorado City	a/ 602.0	b/ 41,443.4	25.1	1.326	.568
3	08123600	Baile Creek above Big Springs	a/ 486.0	b/ 9,001.6	5.3	.347	.149
4	08123800	Baile Creek near Westmore	a/ 959.0	b/ 43,822.0	26.0	.886	.367
5	08123800	Colorado River above Silver	a/ 2,161.0	b/ 89,802.0	50.8	.744	.319
6	08124000	Colorado River at Robert Lake	a/ 33.0	b/ 0.748.0	1.6	1.529	.612
7	08126000	Colorado River at Ballinger	a/ 803.0	b/ 60,184.0	35.7	1.248	.532
8	08127000	Elm Creek at Ballinger	450.0	c/ 54,268.6	32.2	2.280	.969
9	08128000	South Concho River at Christoval	344.0	c/104,916.0	62.2	5.716	2.450
10	08128400	Middle Concho River above Tamerley	1,644.0	93,529.0	19.9	.382	.164
11	08129300	Spring Creek above Tamerley	405.0	a/ 44,200.0	26.2	2.045	.876
12	08130500	Dove Creek at Knickerbocker	218.0	a/ 69,830.0	41.4	6.003	2.573
13	08131400	Pecos Creek near San Angelo	81.1	5,312.0	3.7	1.459	.625
14	08132000	North Concho River at Sterling City	a/ 3,234.0	b/ 5,146.6	1.9	.104	.046
15	08134000	North Concho River near Carlsbad	1,191.0	a/ 6,291.0	3.7	.099	.042
16	08135000	North Concho River at San Angelo	a/ 20.0	b/ 3,245.9	1.9	3.045	1.305
17	08136000	Concho River at San Angelo	4,411.0	b/ 153,660.4	91.0	.653	.280
18	08136000	Concho River at Paint Rock	a/ 1,109.0	b/ 217,492.0	128.9	2.676	1.075
19	08136000	Colorado River near Stacy	a/ 3,234.0	b/ 394,380.0	235.6	2.384	.978
20	08138000	Colorado River at Winnell	a/ 4,231.0	b/ 448,930.0	266.0	1.993	.850
21	08140700	Pecos Bayou near Cross Cut	532.0	a/ 26,549.0	15.7	.935	.401
22	08140800	Jim Ned Creek near Coleman	333.0	b/ 16,142.0	9.6	.908	.389
23	08141500	Honda Creek near Valera	84.2	a/ 27,171.0	1.6	.701	.300
24	08143500	Pecos Bayou at Brownwood	1,660.0	a/ 109,906.0	66.1	12.407	5.317
25	08143600	Pecos Bayou near Mullin	2,073.0	b/ 135,952.0	80.5	1.229	.527
26	08144500	San Saba River at Menard	1,128.0	a/ 145,744.0	86.3	2.421	1.038
27	08144800	Brady Creek near Eden	101.0	a/ 2,610.0	1.6	.484	.207
28	08145000	San Saba River at San Saba	3,038.0	b/ 32,006.3	15.0	.701	.300
29	08146000	San Saba River at San Saba	3,039.0	b/ 381,570.0	126.1	2.353	1.008
30	08147000	Colorado River near San Saba	a/ 9,379.0	b/ 1,106,810.0	655.7	2.213	.948
31	08148500	North Llano River near Junction	814.0	a/ 167,440.0	99.2	3.433	1.471
32	08150000	Llano River near Junction	1,849.0	a/ 372,970.0	221.0	3.780	1.620
33	08150000	Llano River near Mason	3,342.0	a/ 537,140.0	318.2	5.012	1.391
34	08150000	Beaver Creek near Mason	215.0	28,861.0	17.7	2.603	1.116
35	08151500	Llano River at Llano	4,192.0	a/ 749,930.0	444.3	3.353	1.437
36	08152000	Sandy Creek near Kingsland	346.0	138,521.0	82.1	7.503	3.216
37	08153500	Pedernales River near Johnson City	901.0	a/ 470,080.0	278.5	9.777	4.190
38	08158000	Colorado River at Austin	a/ 17,146.0	b/ 3,604,960.0	2,135.5	3.937	1.687
39	08158000	Walnut Creek at Webberville Rd., Austin	51.3	48,773.0	28.9	17.817	7.636
40	08159150	Wilbarger Creek near Pflugerville	4.6	3,894.9	3.3	15.833	6.786
41	08163000	North Fork Guadalupe River near Hunt	169.0	a/ 56,654.0	22.6	6.321	2.709
42	08165000	Guadalupe River at Hunt	288.0	a/ 100,170.0	59.3	6.519	2.784
43	08166000	Johnson Creek near Ingram	114.0	a/ 51,530.0	30.5	8.473	3.632
44	08167000	Guadalupe River at Comfort	839.0	a/ 442,900.0	274.2	10.364	4.438
45	08167500	Guadalupe River near Spring Branch	1,315.0	a/ 902,320.0	534.6	12.862	5.513
46	08167600	Rebecca Creek near Spring Branch	10.9	10,987.0	9.5	18.894	8.098
47	08167800	Guadalupe River at Settler	1,426.0	b/ 989,330.0	586.1	12.914	5.535
48	08168000	Guadalupe Rl. above Comal Rl. at New Braunfels	1,518.0	b/ 1,210,220.0	717.0	14.944	6.408
49	08171000	Blanco River at Wimberly	355.0	a/ 400,980.0	237.6	21.172	9.075
50	08171300	Blanco River near Kyle	412.0	a/ 399,700.0	236.8	18.184	7.794
51	08172000	San Marcos River at Luling	898.0	1,161,230.0	688.0	25.974	1.133
52	08172400	Pium Creek at Lockhart	172.0	a/ 189,396.0	112.3	21.697	9.586
53	08173000	Pium Creek near Luling	309.0	505,150.0	299.3	30.643	3.135
54	08177700	Omos Creek at Oresden Dr., San Antonio	21.2	7,926.0	4.7	7.008	.304
55	08178700	Salado Creek (upper station) at San Antonio	137.0	13,172.0	7.8	1.802	.772
56	08178800	Salado Creek (lower station) at San Antonio	.189.0	a/ 92,960.0	55.1	9.219	3.952
57	08179000	Medina River near Pipe Creek	474.0	a/ 389,640.0	230.8	15.408	6.604
58	08179100	Red Bluff Creek near Pipe Creek	56.3	35,050.0	20.8	11.689	5.022
59	08180800	Medina River near Somerset	967.0	b/ 588,871.7	348.9	11.414	4.892
60	08181400	Medina Creek at Helotes Texas	15.0	a/ 9,001.3	9.3	11.248	4.821
61	08181500	Medina River at San Antonio	1,317.0	b/ 696,460.0	412.6	9.912	4.249
62	08182000	Nueces River at Laguna	764.0	a/ 312,080.0	184.9	7.657	3.281
63	08190500	West Nueces River near Brackettville	700.0	a/ 55,291.4	32.8	1.481	.635
64	08192000	Nueces River below Uvalde	1,967.0	a/ 359,610.0	213.0	3.462	1.484
65	08195000	Nueces River at Concan	405.0	a/ 351,440.0	149.0	11.637	4.897
66	08196000	Dry Frio River near Reagan Wells	117.0	a/ 82,870.0	49.1	13.276	5.860
67	08197000	Frio River below Dry Frio River near Uvalde	266.0	a/ 32,697.5	89.4	.927	.393
68	08198000	Sabinal River near Sabinal	601.0	a/ 150,840.0	89.4	7.372	3.082
69	08199000	San Antonio River at San Antonio	547.0	a/ 87,360.0	184.0	6.753	2.817
70	08200000	Hondo Creek near Tarpley	86.2	a/ 106,920.0	63.3	23.280	9.864
71	08202700	Hondo Creek at King Metehole near Hondo	144.0	a/ 12,095.9	7.2	1.997	.884
72	08201500	Solo Creek at Miller Ranch near Utopia	168.0	a/ 36,884.0	21.9	16.041	6.875
73	08202700	Solo Creek at Rose Ranch near D'Arms	168.0	f/ 4,021.4	2.4	.449	.192
74	08276300	Sanderson Canyon at Sanderson	195.0	3,319.0	2.0	.319	.137
75	08412500	Pecos River near Fort	21,210.0	b/ 144,399.0	85.5	1.28	.055
76	08431700	Limpia Creek above Fort Davis	52.4	2,101.3	1.2	.752	.322
77	08431700	Limpia Creek below Fort Davis	227.0	7,238.0	4.3	.698	.256
78	08428900	Guadalupe River at Fort Stockton	1,034.0	a/ 18,000.0	1.0	.016	.007
79	08446500	Pecos River near Gilvin	29,560.0	b/ 185,311.5	109.8	11.8	.518