

HYDROGEOLOGY OF THE MIDDLE WILCOX AQUIFER  
SYSTEM IN MISSISSIPPI

INTRODUCTION

This report describes the hydrogeology of the middle Wilcox aquifer system in Mississippi and its present use and potential for future freshwater development. The purpose of this report is to provide water-planners and managers with hydrogeologic data on this predominantly undeveloped aquifer system and to describe in potential as an alternative source of water in the State. This report was prepared to complete a series of maps describing the hydrogeology of the major aquifers in Mississippi.

This report was prepared as part of the Gulf Coast Regional Aquifer System Analysis (RASA) project using hydrogeologic data assembled for the flow analysis of aquifers in Tertiary sediments in the Mississippi embayment study area (Arthur and Taylor, 1989).

LOCATION, STRATIGRAPHY, AND THICKNESS

Sediments of the Wilcox Group in Mississippi are exposed in a crescent-shaped belt extending from Benton and Tippah Counties in the north to Clarke, Kemper, and Lauderdale Counties in the east (fig. 1). The subsurface beds dip westward in the northern and western parts of the State because of influence of the Mississippi embayment trough, which roughly parallels the Mississippi River. In the southern part of the study area, the direction of dip gradually shifts to the southwest toward the Gulf Coast geocline. The rate of dip ranges from about 20 feet per mile in the northern part of the study area to more than 50 feet per mile in the southern part. In the Jackson area, the trend of the dip is interrupted by the Jackson dome, a local structural uplift (fig. 1).

In Mississippi, the Wilcox Group of Paleocene and Eocene ages includes, in ascending order, the Natchez, Tuscaloosa, and Hachisette Formations (table 1). The Wilcox Group generally is undifferentiated throughout most of Mississippi. However, in areas where the Wilcox Group is thicker, the beds are divided into an important regional aquifer—the Meridian-upper Wilcox aquifer—made of the upper part of the Wilcox Group and the overlying Meridian Sand Member of the Tallahatchie Formation (Boswell, 1976). The lower part of the Wilcox Group and the uppermost part of the underlying Midway Group in this area are also marked by thick sand beds that make up the important regional aquifer, the lower Wilcox aquifer (Boswell, 1976). In areas where the Wilcox Group is thinner, the beds are often irregular, discontinuous, and of local extent; however, the sand beds function as a single hydraulic unit and are collectively referred to as the middle Wilcox aquifer system and may represent a significant source of ground water. The intervening confining units consist mainly of clay. For a more detailed description of the units, the reader is referred to the report by Arthur and Taylor (1989).

The thickness of the middle Wilcox aquifer system increases from a few hundred feet in the outcrop area to more than 1,000 feet in the extreme down-dip sections of the study area (sheet 2). Downward from the outcrop area, total accumulated thickness of sand beds that are at least 20 feet thick ranges from less than 100 feet in the northern part of the study area to more than 400 feet in three localities in the central and south-central part of the study area (fig. 2). At least one sand bed 100 feet or more thick is present in a 15-county area that extends from Rankin and Smith Counties in the south to Tate County in the north.

GROUND WATER

The principal source of recharge to the middle Wilcox aquifer system is from precipitation in the outcrop area (fig. 1). Precipitation in the area ranges from about 50 to 56 inches per year, but of that amount about 14 to 20 inches per year leaves the area as surface runoff. Most of the water that percolates into the ground is lost by evapotranspiration or ground-water discharge to local streams.

Regional movement of ground water in the middle Wilcox aquifer system is from areas of high potentiometric levels to areas of low potentiometric levels—generally westward except in the southern part of the study area where movement is to the south-southwest. Locally, the rate and direction of ground-water movement are controlled by the hydraulic conductivity of the sand bed and by the hydraulic conductivity of the confining units.

The potentiometric surface shown in figure 3 represents the approximate altitude of water levels in wells screened in the middle Wilcox aquifer system in 1985. Although the contours were based on water-level measurements throughout the year, the contours are considered representative because seasonal water-level changes are generally small (less than 5 feet) and because of the large contour interval (50 feet). Because the potentiometric map is based on water-level measurements from wells screened in different sand beds within the aquifer system, the water level of an individual sand bed could differ from that shown.

Near some pumping centers in and near the recharge area, water-level declines in recent years have been in the range of about 0.5 to 1.0 foot per year (fig. 4). Generally, however, the middle Wilcox aquifer system is unmineralized (Arthur and Taylor, 1989) and no significant cones of depression occur in the aquifer.

WATER WELLS AND AQUIFER CHARACTERISTICS

Water wells have been completed in sand beds of the middle Wilcox aquifer system in 23 counties in Mississippi. Most of the several thousand water wells in this aquifer system are small-diameter domestic wells located in or immediately down-dip from the outcrop area. Although yields from most of these wells are small, the aquifer system is capable of yielding large quantities of water to larger wells. More than 90 wells in the middle Wilcox aquifer system yielded more than 100 gallons per minute; some wells yielded as much as 500 gallons per minute.

The availability of freshwater at shallower depths has limited the development of this aquifer system farther down-dip. However, some water wells in the aquifer are more than 700 feet deep in the southern part of the study area and more than 800 feet deep in the western and northern parts of the study area. A few wells are more than 1,000 feet deep. More than 100 public and industrial wells in 19 counties in Mississippi obtain water from the middle Wilcox aquifer system.

Results of aquifer tests in the middle Wilcox aquifer system are sparse, but tests have been completed at 7 sites in 5 counties (Newcome, 1971), and specific capacity data are available for 15 sites in 9 counties (fig. 5). Well and aquifer characteristics obtained at these sites are summarized in the following table.

	Transmissivity (feet squared per foot)	Hydraulic conductivity (feet per day)	Specific capacity (gallons per minute per foot)
Maximum	2,800	46	22
Median	330	16	3
Minimum	260	4	0.5
No. of tests	7	7	15

Although aquifer characteristics in the middle Wilcox aquifer system are variable, results from aquifer tests can be used to estimate water-level drawdowns in pumped wells and to plan well-spacing to minimize well interference. If transmissivity values of the aquifers and pumping rates of wells are known or can be estimated, evaluations of water-level drawdown during pumping can be made from figure 6. A well that is being pumped influences the water levels in nearby wells. The extent of the influence is a function of the rate of discharge, the transmissivity of the aquifer, and the time that has elapsed since pumping began. The influence that pumping a well will have on other wells in the same aquifer may be determined if the transmissivity and coefficient of storage of an aquifer are known or can be estimated (fig. 7). In planning well-spacing and withdrawal rates, consideration of drawdown effects for various combinations of time and distance is necessary.

GROUND-WATER QUALITY

Ground water in the outcrop area of the middle Wilcox aquifer system generally is of a mixed, calcium sodium bicarbonate type. Dissolved-solids concentrations generally are very small in the outcrop area but increase down-dip to the west and southwest. As the water becomes more mineralized with increased distance from the outcrop area, sodium and bicarbonate become the major constituents in the water. On the basis of interpretation of geophysical log data, primarily during the Gulf Coast RASA investigation (Arthur and Taylor, 1989), freshwater—water with a dissolved-solids concentration of less than 1,000 milligrams per liter—extends about 70 miles from the outcrop in central Mississippi but only extends about 20 miles from the outcrop area in extreme northern and east-central Mississippi (fig. 1). Freshwater may be obtained from the middle Wilcox aquifer system in all or parts of 31 counties in Mississippi.

There is a general trend of increasing pH values and concentrations of dissolved sodium, bicarbonate, nitrate, and iron with increasing depth for the aquifer system. Typically, water in the middle Wilcox aquifer system has concentrations much smaller than the recommended limits for drinking water for nitrate (10 milligrams per liter as nitrogen), sulfate (250 milligrams per liter), and fluoride (2 milligrams per liter) (U.S. Environmental Protection Agency, 1966a, b). Excessive concentrations of iron and color are common water-quality problems in parts of the lower Wilcox and the Meridian-upper Wilcox aquifers (Boswell, 1975, 1976), and it seems logical to assume they may be problems in some parts of the middle Wilcox aquifer system. Although iron concentrations have been determined in water from only a few wells in the middle Wilcox aquifer system, these concentrations did not exceed the recommended limit of 300 micrograms per liter (U.S. Environmental Protection Agency, 1966b) for drinking water. The recommended limit for color in drinking water (15 units) was exceeded in water from about 20 percent of the wells in the middle Wilcox aquifer system; the color of water from one well in Lauderdale County was 30 units (table 2).

Temperature of shallow ground water (about 100 feet or less deep) ranges from about 15.5 °C (60 °F) in Benton County to about 18.5 °C (65 °F) in Lauderdale County, and as an annual basis, approximates the mean annual air temperature. Ground-water temperature increases with depth at a rate of about 2.5 °C per 100 meters or about 1.4 °F per 100 feet (fig. 8).

WATER USE

In 1985, the middle Wilcox aquifer system was the source of water for about 38 public-supply systems and 7 large industries in Mississippi. Data are from the U.S. Geological Survey Mississippi Site-specific Water Use Data System (MWDUS). Total withdrawal of water for these uses was about 10 Mgal/d (million gallons per day) in 1985 (fig. 9 and table 3). About 40 Mgal/d were withdrawn for public supply and about 3.9 Mgal/d were used for irrigation. The largest withdrawals were in Grenada County where 5.2 Mgal/d were withdrawn primarily for irrigation and industry. The largest public-supply system was the town of Adamsville in Choctaw County, which withdrew about 1.6 Mgal/d in 1985. Although thousands of small wells obtain water from this aquifer system for domestic and other uses, the total withdrawal for these uses is probably less than 1 Mgal/d.

POTENTIAL FOR FUTURE DEVELOPMENT

Regionally, the middle Wilcox aquifer system is a relatively undeveloped source of ground water in Mississippi. Although well yields in excess of 100 gallons per minute are possible, pumping from this aquifer system is less than from any other of the extensive aquifers in the State. For this reason, water levels have not declined in this aquifer system as they have in other large aquifer systems. Water from this aquifer system generally is suitable for most uses.

The middle Wilcox aquifer system has the capability of serving as a source of water for much of Mississippi. Water from this aquifer system could augment existing water-supply systems locally if the adjacent lower Wilcox and Meridian-upper Wilcox aquifers cannot supply the demand for water or contain water of unsuitable quality.

REFERENCES

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- Dalsin, G.L., 1978. Water for industrial and agricultural development in Bolivar, Carroll, Leflore, Sunflower, and Tallahatchie Counties, Mississippi: Mississippi Research and Development Center Bulletin, 67 p.
- Newcome, Roy Jr., 1971. Results of aquifer tests in Mississippi: Mississippi Board of Water Commissioners Bulletin 71-2, 44 p.
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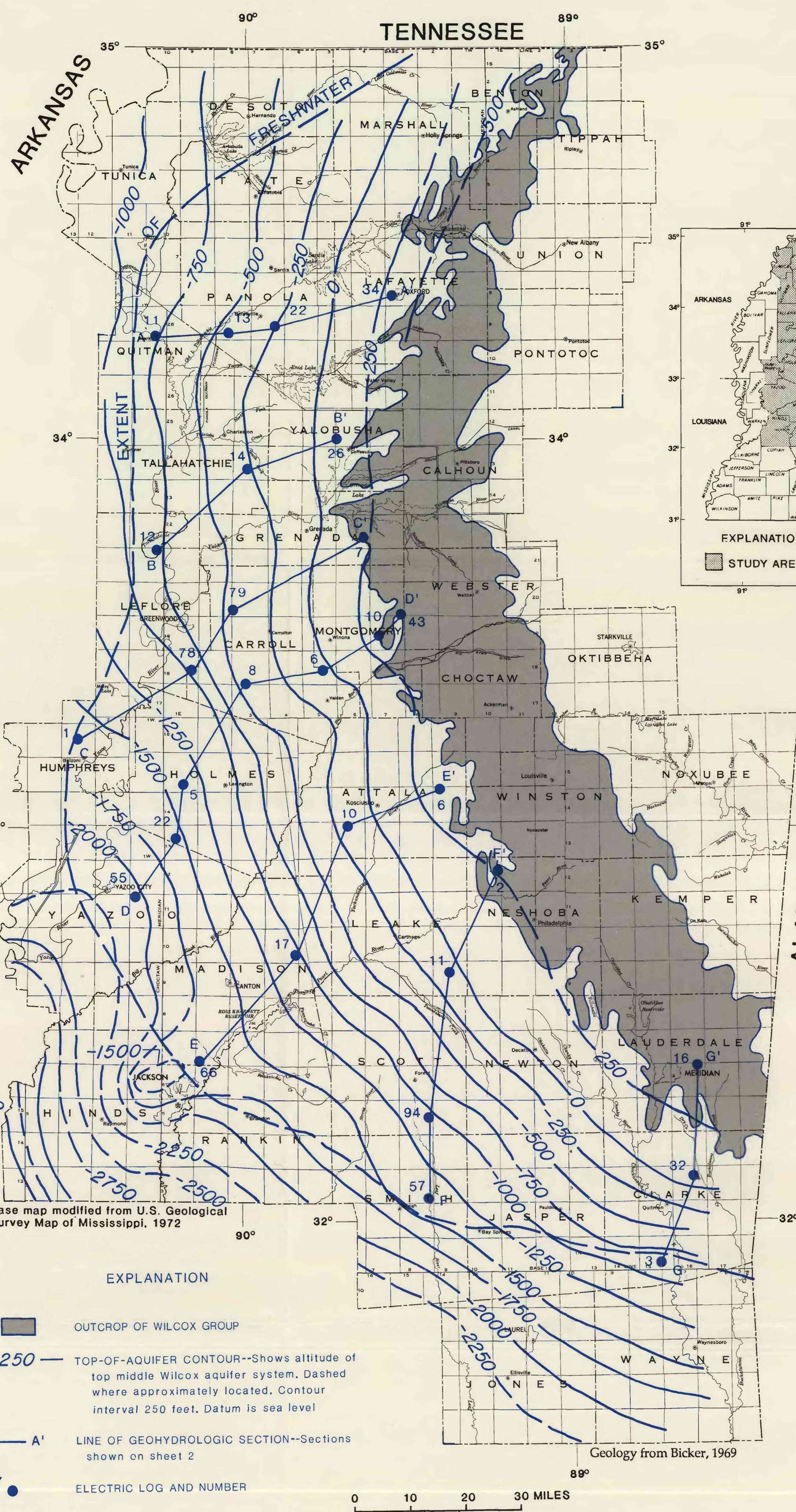


Figure 1.—Location of study area, outcrop of the Wilcox Group, configuration of the top of the middle Wilcox aquifer system, and locations of geophysical sections.

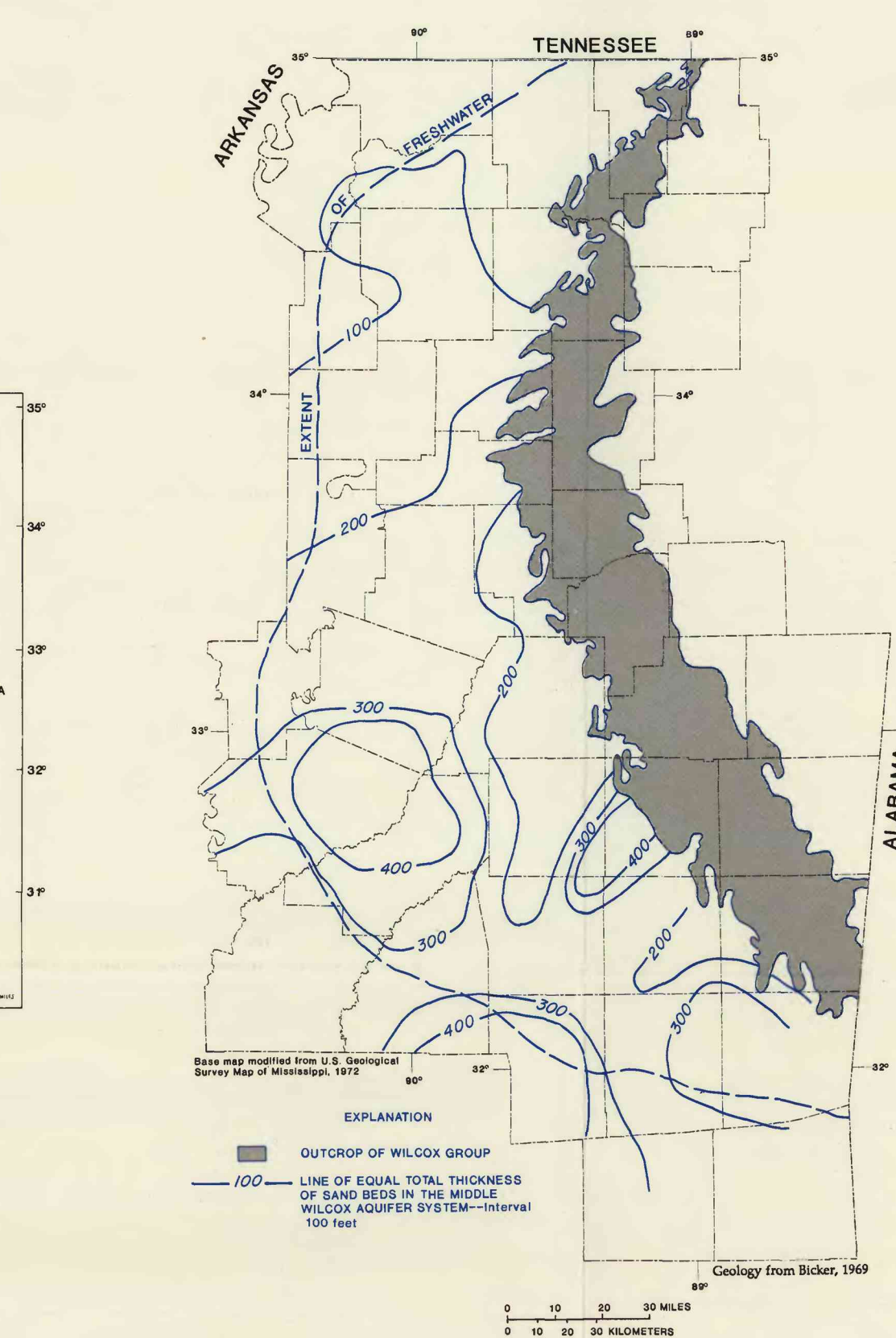


Figure 2.—Total thickness of sand beds in the middle Wilcox aquifer system.

FACTORS FOR CONVERTING INCH-POUND UNITS TO INTERNATIONAL SYSTEM (SI) UNITS	
Multiple inch-pound units	To obtain metric units
Foot (ft)	meter
mile (mi)	kilometer
square mile (sq mi)	square kilometer
million gallons per day (Mgal/d)	cubic meters per second
foot per mile (ft/mi)	centimeter per kilometer
foot squared per day (ft²/d)	centimeter squared per day
gallons per minute per foot (gpm/ft)	liter per second per meter
ft²/mi²	ft²/mi²

Throughout this report water temperatures are reported in degrees Celsius (°C). Temperatures may be converted to degrees Fahrenheit (°F) equivalent with the following formula:

$$^{\circ}\text{F} = 1.8^{\circ}\text{C} + 32$$

Sea Level: In this report "sea level" refers to the National Geodetic Vertical Datum of 1955 (NGVD of 1955)—a geoid datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called Sea Level Datum of 1929.

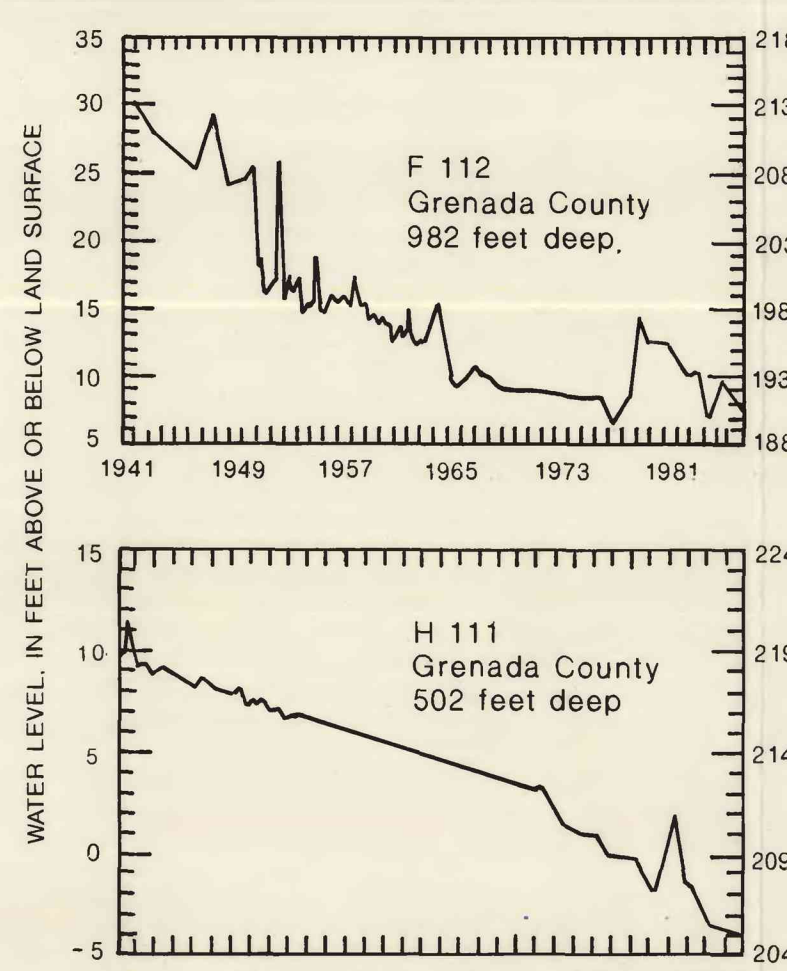


Figure 4.—Water-level trends in observation wells in the middle Wilcox aquifer system.

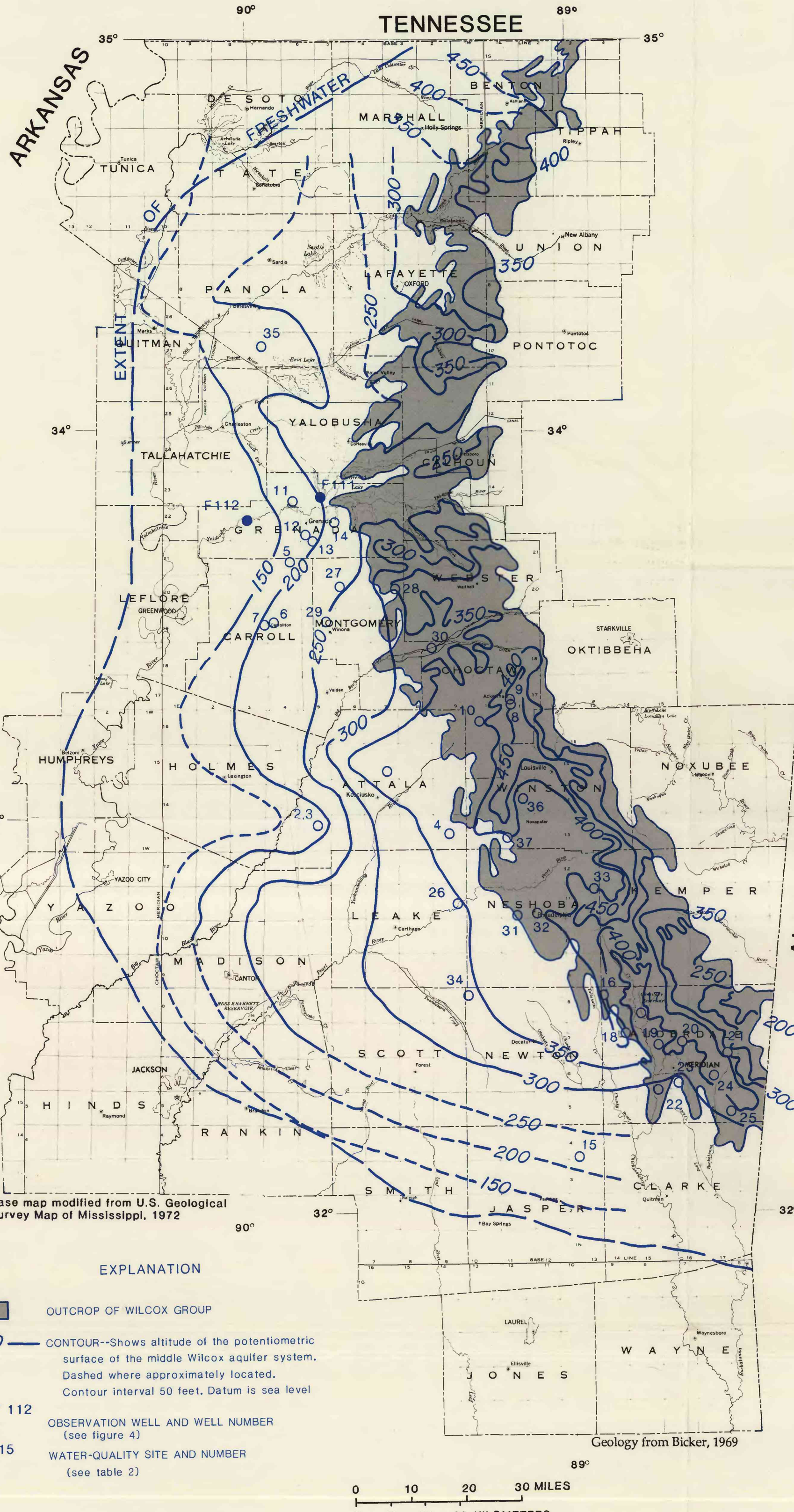


Figure 3.—Potentiometric surface of the middle Wilcox aquifer system.

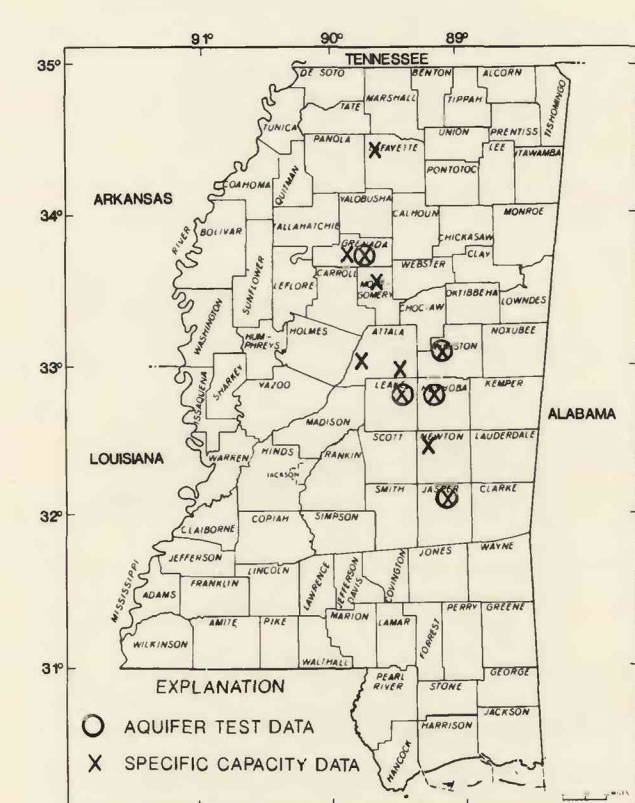


Figure 5.—Location of well sites with aquifer and specific capacity data.

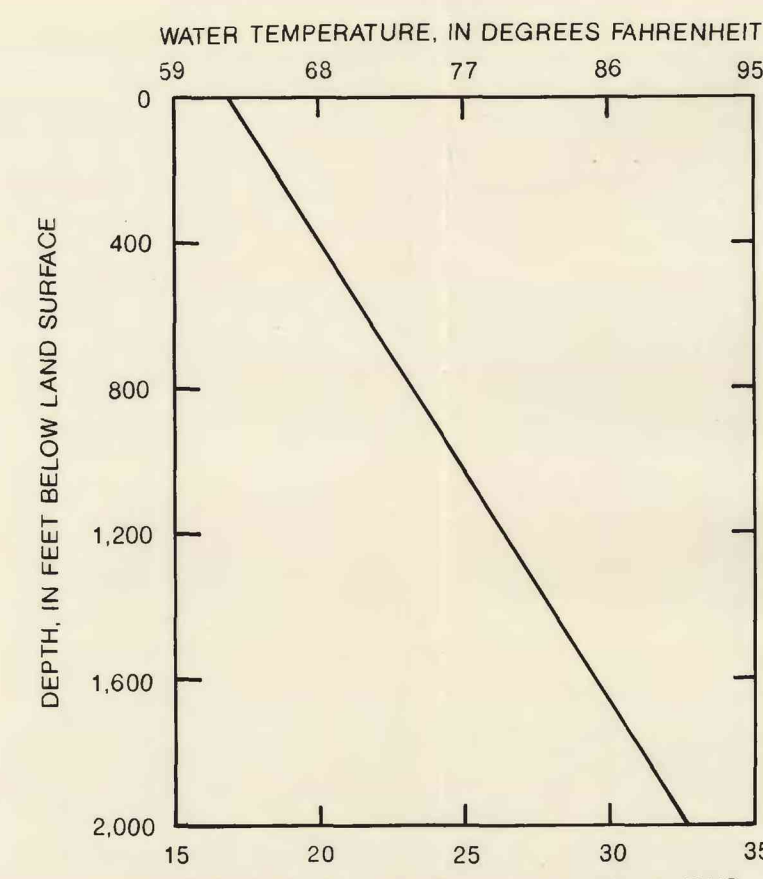


Figure 8.—Typical depth-temperature relation for the middle Wilcox aquifer system. (Modified from Dalsin, 1978)

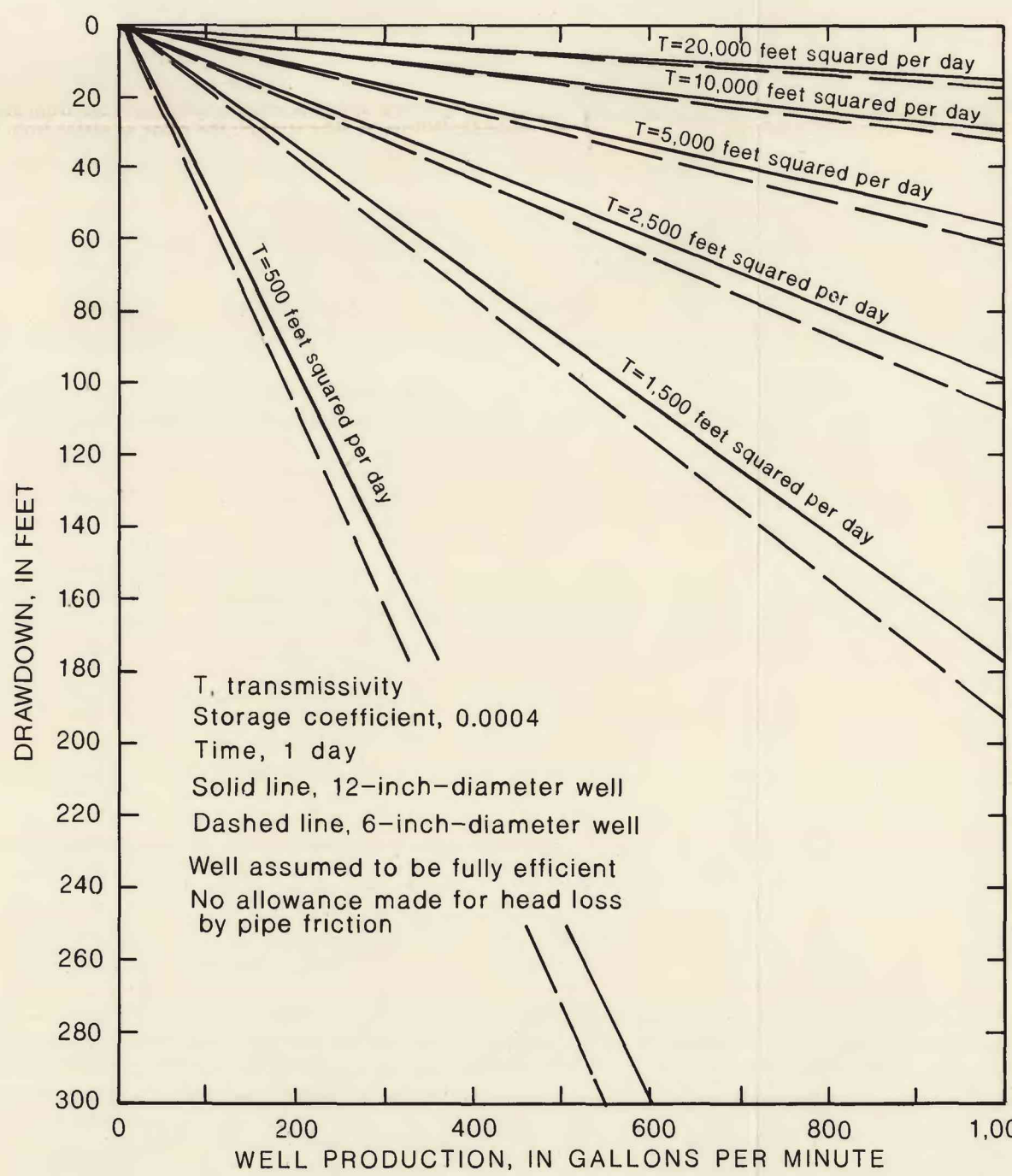


Figure 6.—Drawdown to be expected in pumped wells tapping an artesian aquifer of the study area (modified from Dalsin, 1978).

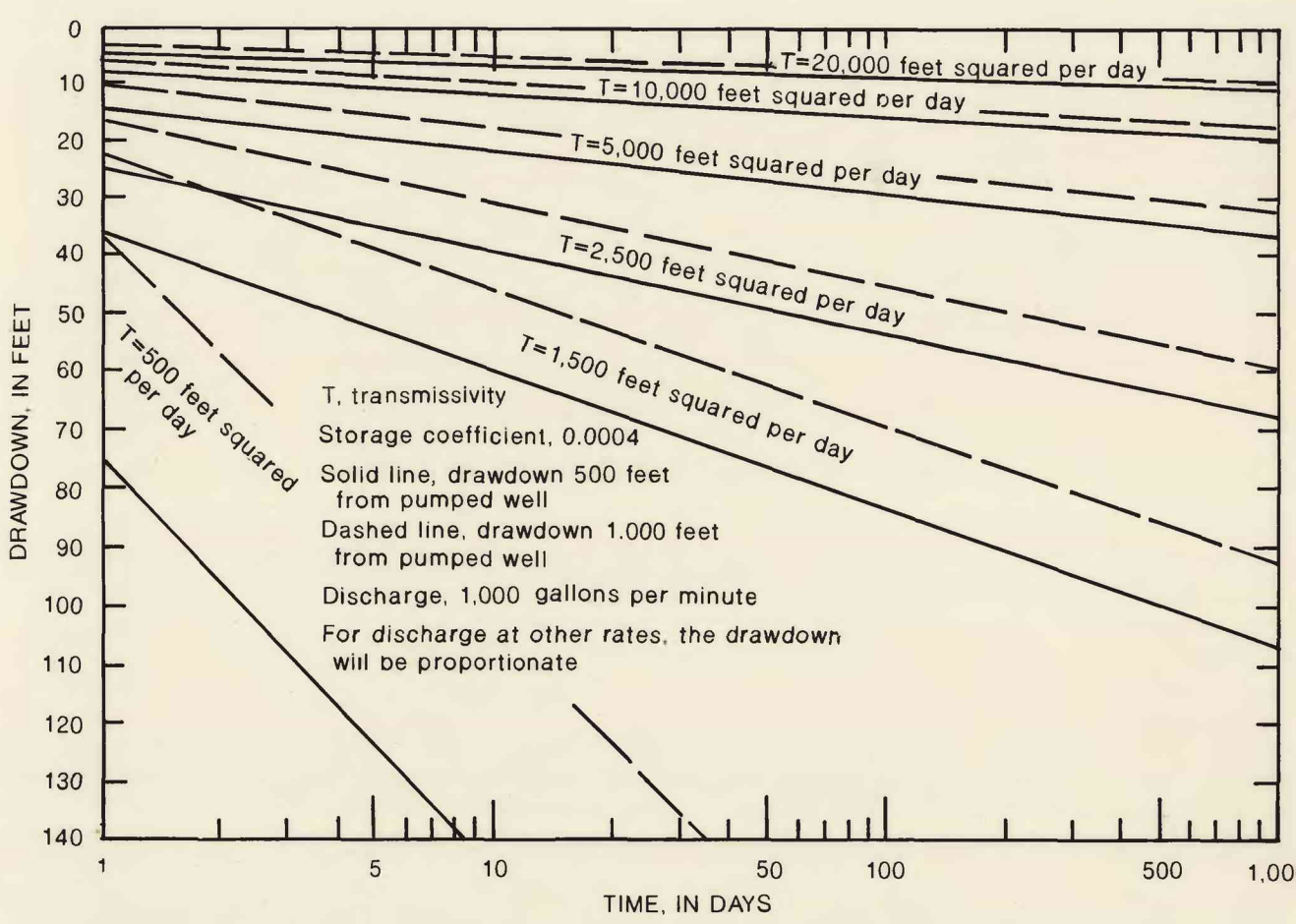


Figure 7.—Theoretical time-distance-drawdown relations for pumping from an artesian aquifer of the study area (modified from Dalsin, 1978).

Table 2.—Physical and chemical characteristics of water from selected wells in the middle Wilcox aquifer system in Mississippi

[Units in milligrams per liter, except as noted: ft, foot; g/L, milligrams per liter; dashes indicate missing values]														
[Location]														
Site no.	Depth (ft)	Date (year-month)	pH (unit)	Color (units)	Salinity (ppt)	Calc. (mg/L)	Magnesium (mg/L)	Sodium (mg/L)	Potassium (mg/L)	Bicarbonate (mg/L)	Sulfate (mg/L)	Chloride (mg/L)	Total dissolved solids (mg/L)	Total suspended solids (mg/L)
ATLANTA COUNTY														
1	561	63-11-07	7.3	1	130	19	62	11	40	108	72	28	0.1	0.09
2	1,000	62-07-07	8.2	30	130	12	2	130	8	—	2.0	4.9	1	—
3	1,000	64-01-09	7.1	19	21	19	43	2.2	—	—	15	2.4	1	160
4	520	72-03-29	7.8	5	170	31	5.5	14	34	130	15	7.9	1	22
CARROLL COUNTY														
5	828	85-05-27	8.3	30	280	22	6	110	1.3	—	9	14	2	303
6	727	72-10-19	8.2	5	308	31	5	120	3.0	292	1.0	1.0	2	40
7	549	79-08-28	8.6	5	483	21	4	160	1.5	300	9	17	2	—
CHOCTAW COUNTY														
8	314	79-03-12	5.4	—	41	48	0	97	1.8	—	32	12	0	—
9	112	65-08-29	5.4	—	11	16	1.7	—	—	0	—	—	—	—
10	3	75-01-09	5.4	0	34	30	8	2.9	9	16	4	4.8	3	10
GRENADE COUNTY														
11	633	71-03-10	8.4	20	120	4.4	1.2	162	2.2	234	2	47	2	16
12	489	71-03-10	7.8	10	20	11	1.2	147	1.3	248	2	35	2	—
13	331	69-04-30	8.4	15	240	4.4	1.8	102	5	234	2.4	16	1	14
14	482	71-03-10	7.5	5	240	34	8	18	88	15	234	2.4	16	1
JASPER COUNTY														
15	886	67-05-31	8.1	40	220	2	2	32	8	344	9.0	2.4	2	04
LAUDERDALE COUNTY														
16	185	48-09-26	6.6	30	100	10	2.4	8.4	1.5	56	4.4	2.4	1	02
17	388	48-09-26	7.2	10	160	21	6	37	11	152	10	2.0	0	00
18	200	48-09-27	7.2	10	160	17	1.4	1.8	162	4.8	2.2	0	02	—
19	305	48-09-27	7.4	10	160	17	1.4	1.8	162	4.8	2.2	0	02	—
20	630	48-09-27	6.9	300	140	17	6.0	5.7	2.7	300	0	2.8	2	07
21	270	48-09-27	7.4	10	160	17	1.4	1.8	162	4.8	2.2	0	02	—
22	300	47-05-13	7.8	10	160	5	4	73	5	176	12	2.9	1	00
23	165	47-05-13	7.8	10	160	5	4	73	5	176	12	2.9	1	00
24	630	48-09-19	7.4	5	170	16	5	45	1.2	158	10	2.1	0	02
25	310	48-09-19	7.2	5	170	16	5	45	1.2	158	10	2.1	0	02
LEAKE COUNTY														
26	849	85-07-01	—	5	170	12	1	48	1.3	—	97	1.9	1	18
MONTGOMERY COUNTY														
27	330	85-07-29	—	4	190	2.6	4	72	1.3	—	2.8	5.8	1	10
28	386	79-03-24	8.3	40	180	14	1.5	45	2.2	80	3.2	6.1	1	—
29	386	79-03-24	7.8	5	180	14	1.5	45	2.2	80	3.2	6.1	1	—
30	73	13-04-14	—	—	300	22	18	—	—	—	93	28	30	—
NESHOBA COUNTY														
31	727	79-03-13	8.5	0	115	89	1.9	—	—	29	13	12	—	—
32	312	77-05-19	6.2	—	50	17	1.2	5.1	1.2	21	29	8.9	—	—
NEWTON COUNTY														
34	325	85-05-07	8.3	5	180	1.6	4	64	1.8	—	8.4	2.5	1	—
PANOLA COUNTY														
35	823	72-03-17	7.4	0	160	9.9	2.5	44	3.2	152	5.2	1.1	1	—
WINSTON COUNTY														
36	467	70-05-05	7.8	5	210	47	4.5	18	2.8	218	0	4.3	1	02
37	525	79-01-18	8.1	1	170	30	11	11	210	10	3.8	2.2	0	07

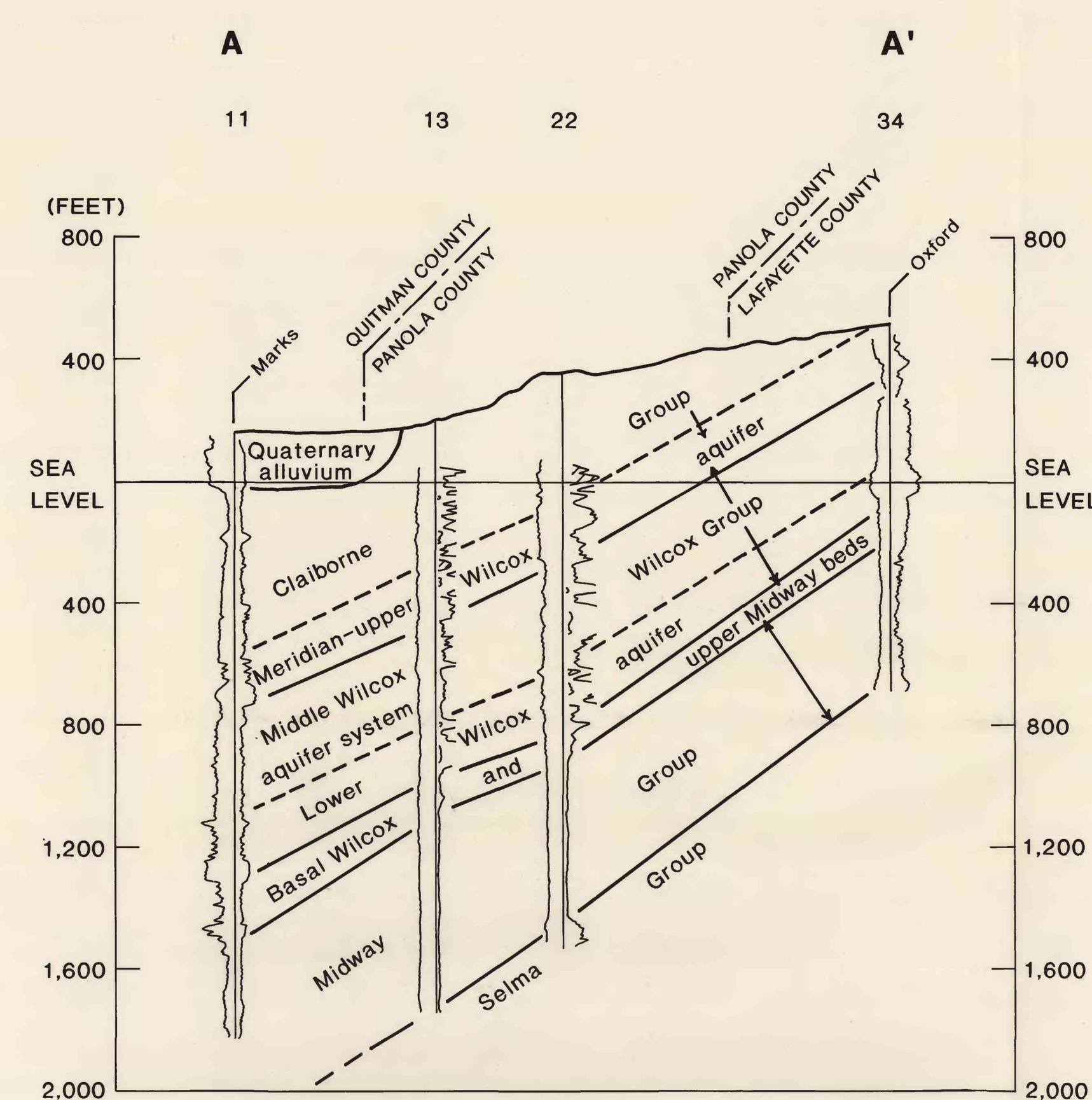


Figure 10.—Geohydrologic section A-A' from Quitman County to Lafayette County, Mississippi.

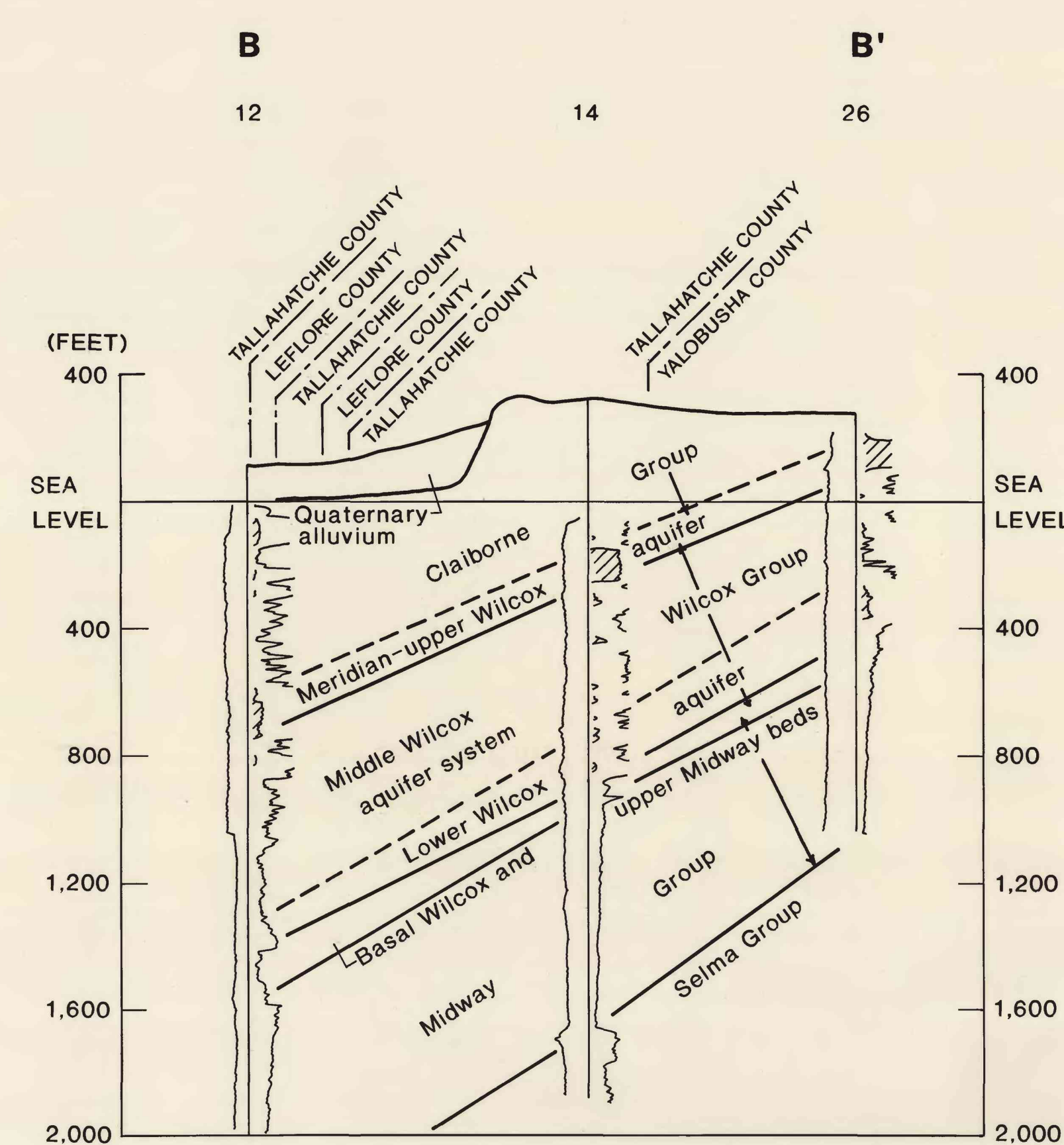


Figure 11.—Geohydrologic section B-B' from Tallahatchie County to Yalobusha County, Mississippi.

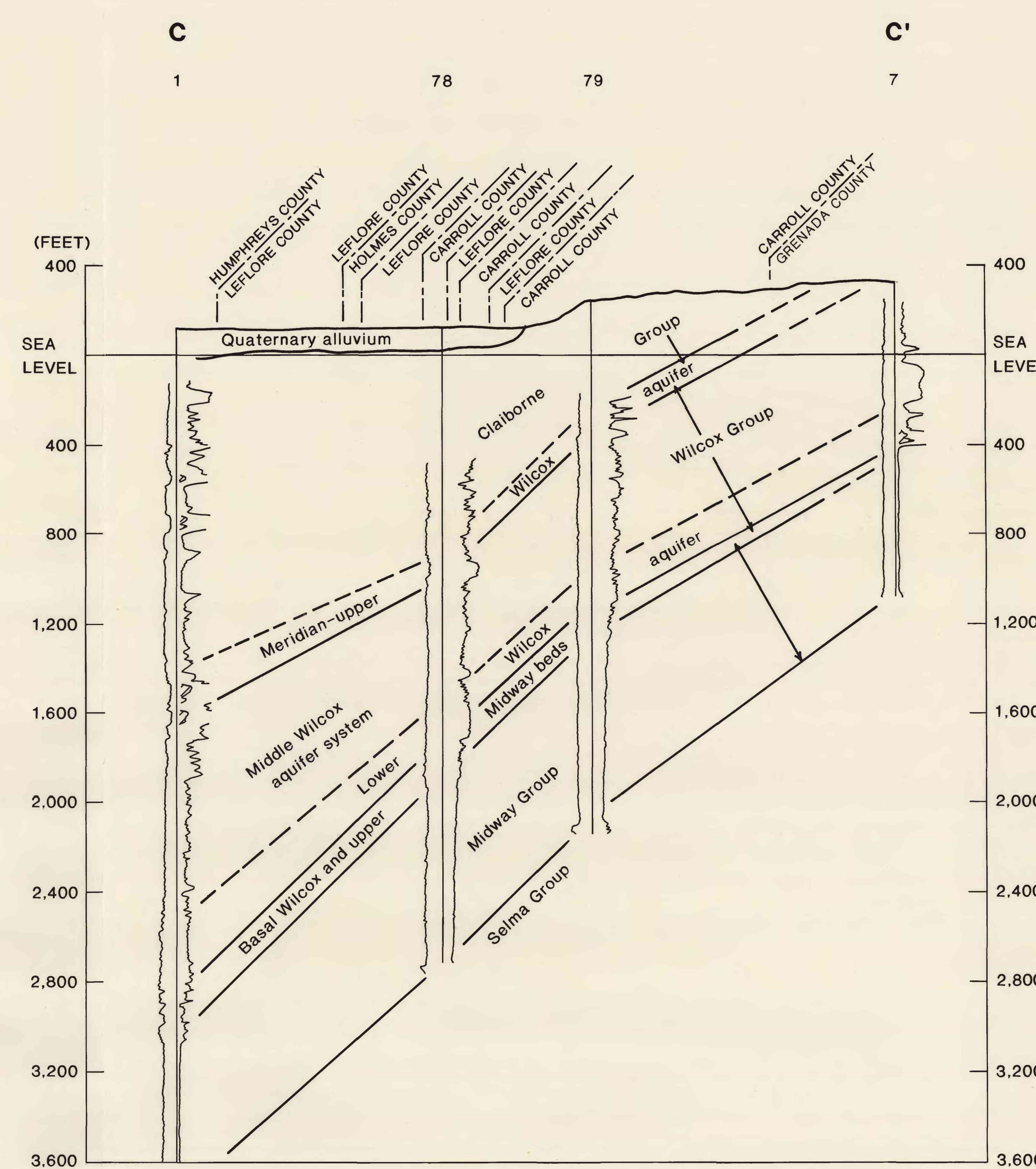


Figure 12.—Geohydrologic section C-C' from Humphreys County to Grenada County, Mississippi.

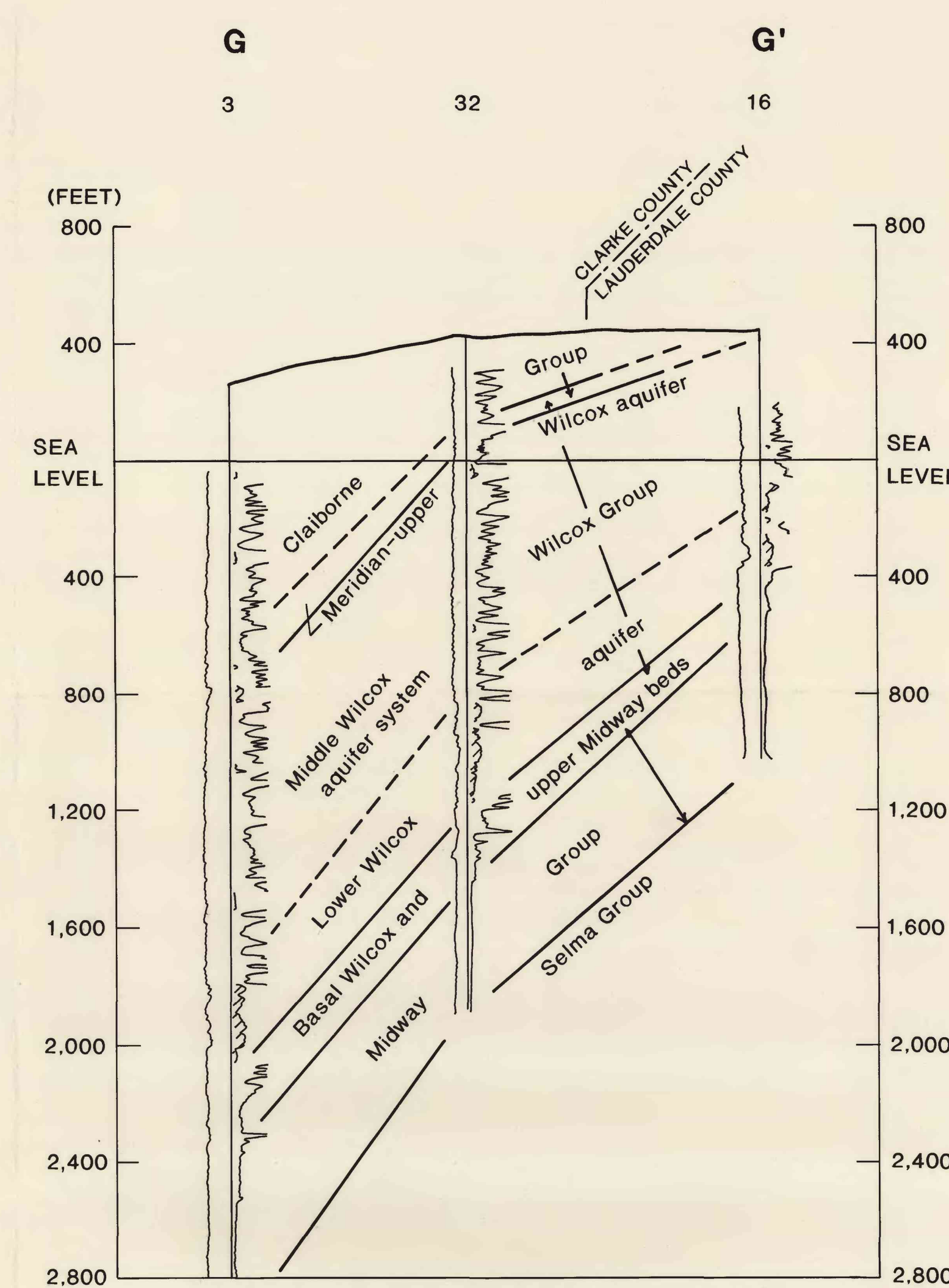
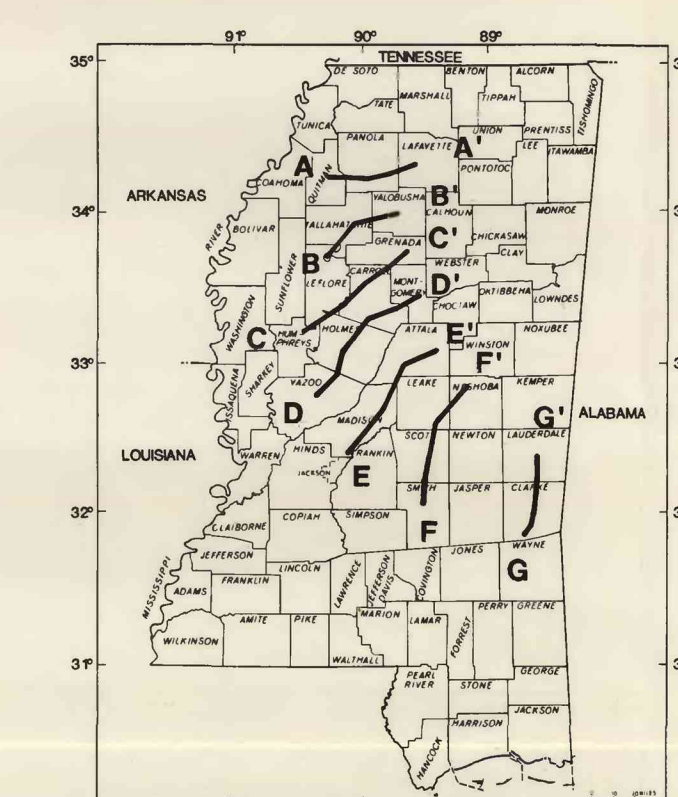


Figure 16.—Geohydrologic section G-G' from Clarke County to Lauderdale County, Mississippi.



GENERALIZED LOCATION OF SECTION LINES

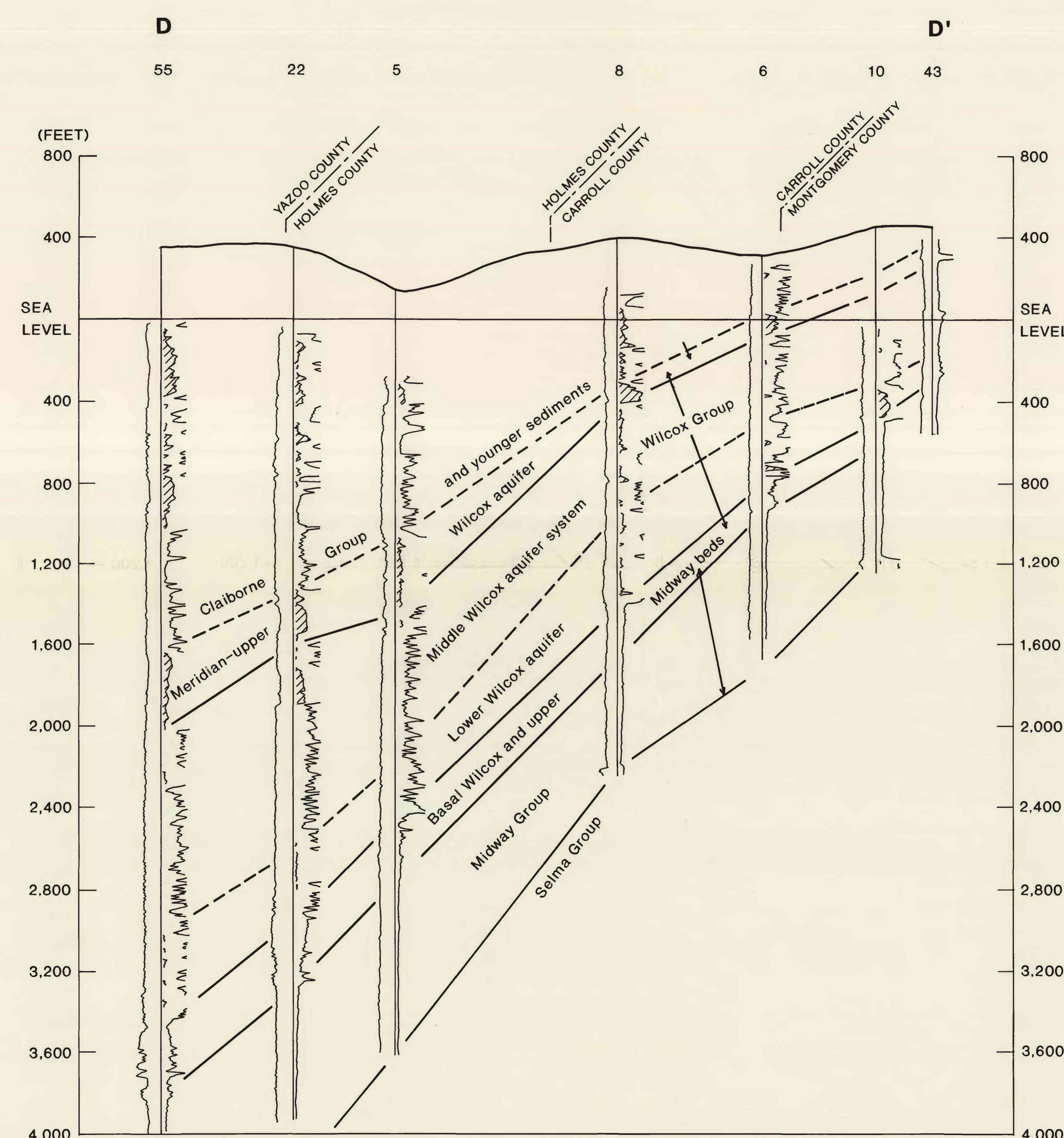
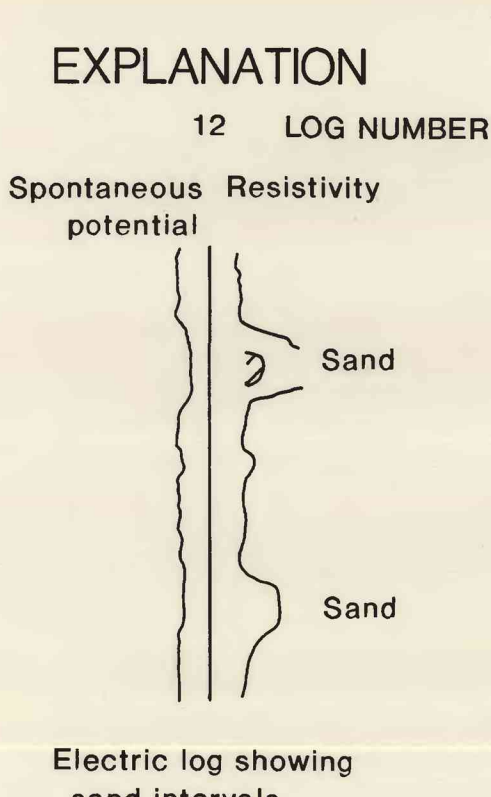


Figure 13.—Geohydrologic section D-D' from Yazoo County to Montgomery County, Mississippi.

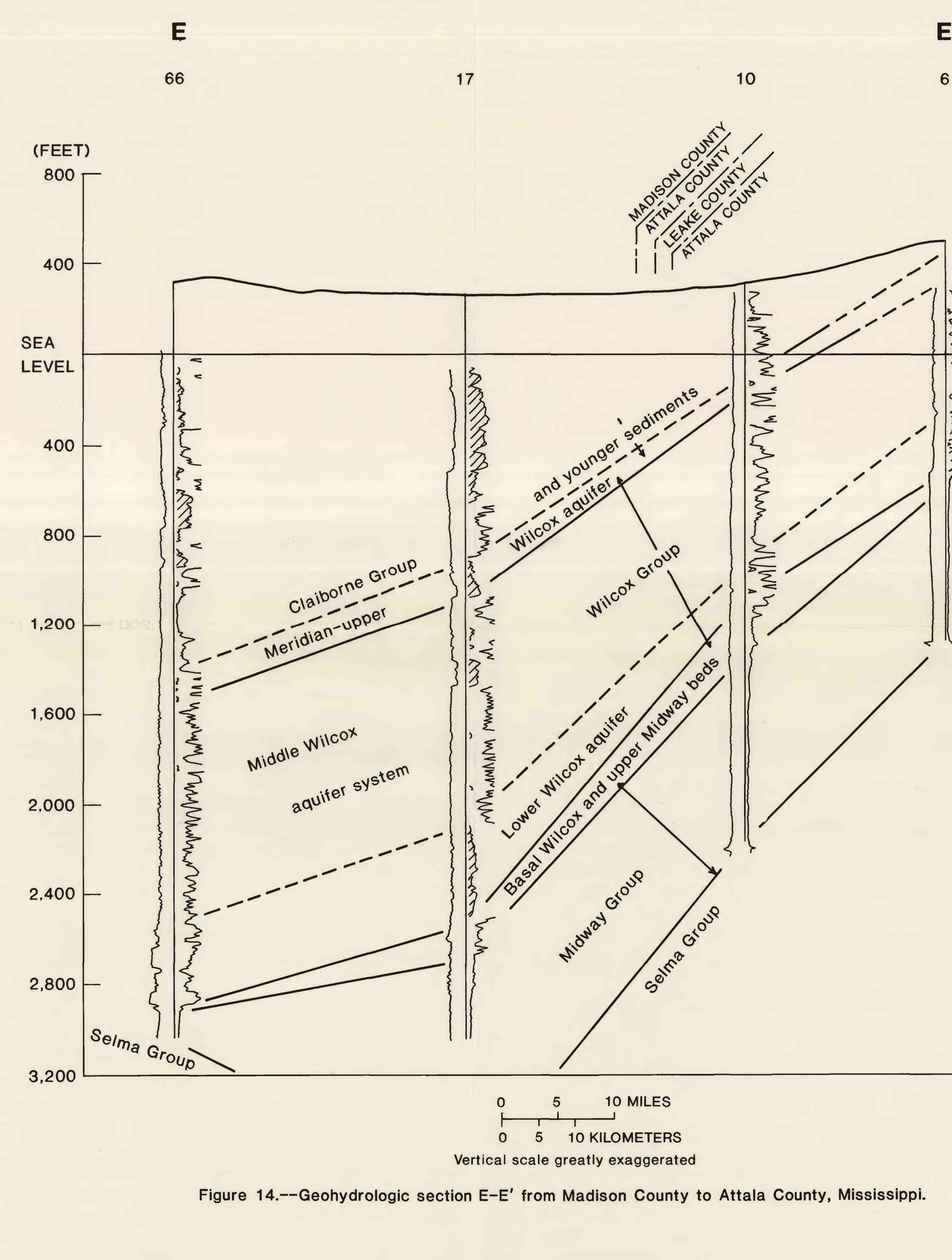


Figure 14.—Geohydrologic section E-E' from Madison County to Attala County, Mississippi.

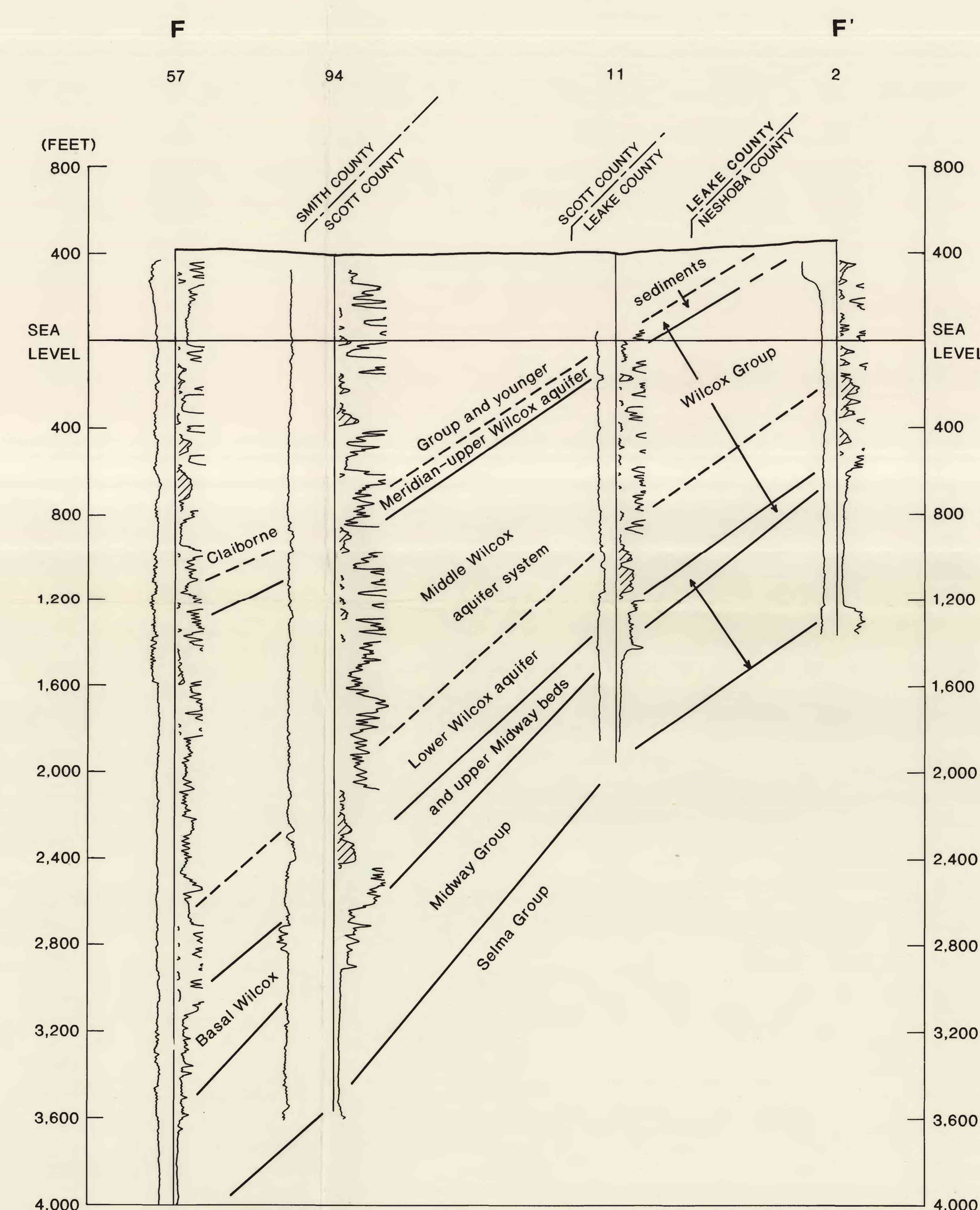


Figure 15.—Geohydrologic section F-F' from Smith County to Neshoba County, Mississippi.

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