

GEOHYDROLOGY OF THE MORRISON FORMATION IN THE WESTERN SAN JUAN BASIN, NEW MEXICO

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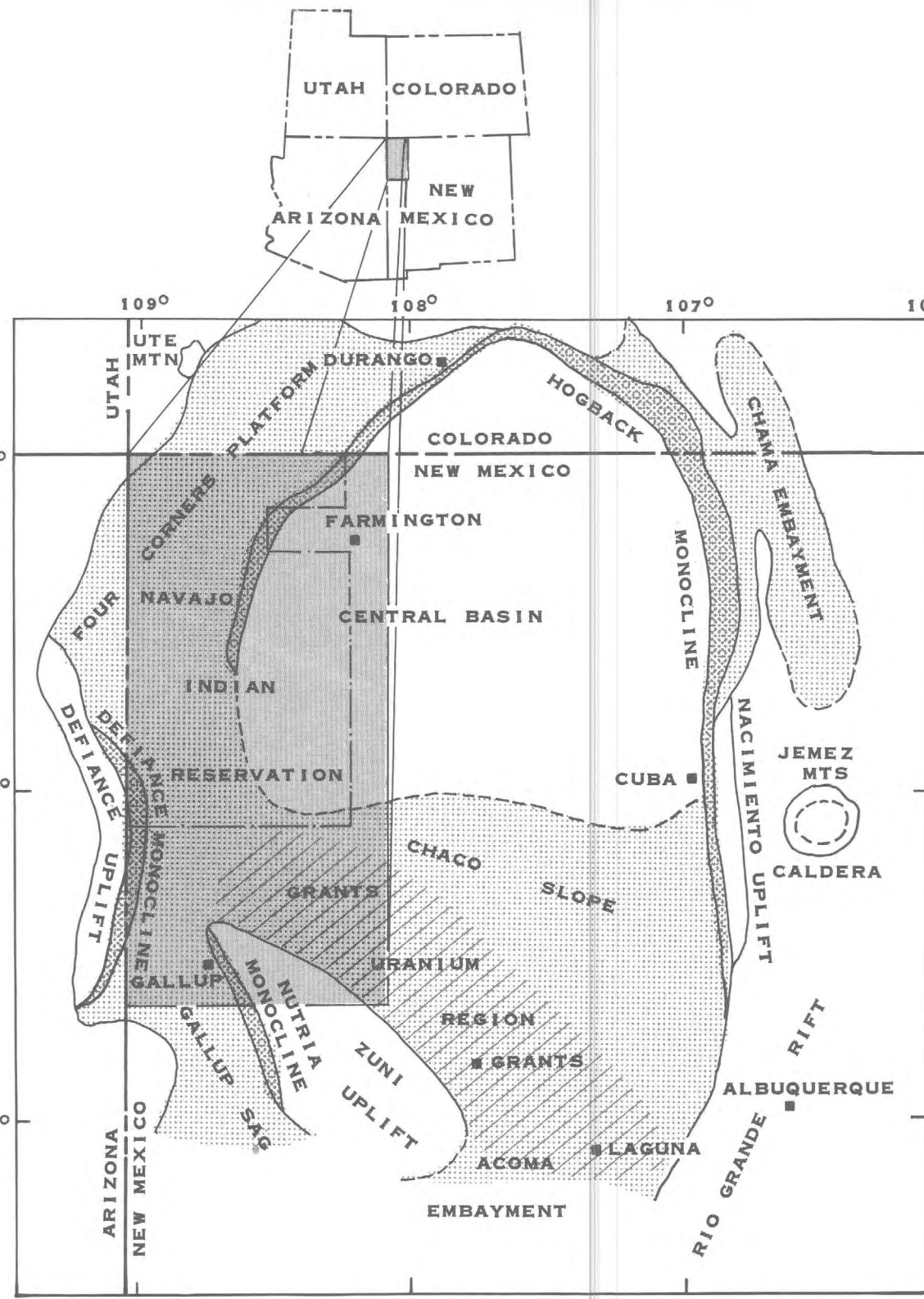


Figure 1.--San Juan Basin showing major structural features, Grants uranium region, and location of the study area (modified from Kelley, 1951; Santos and Turner-Peterson, 1986).

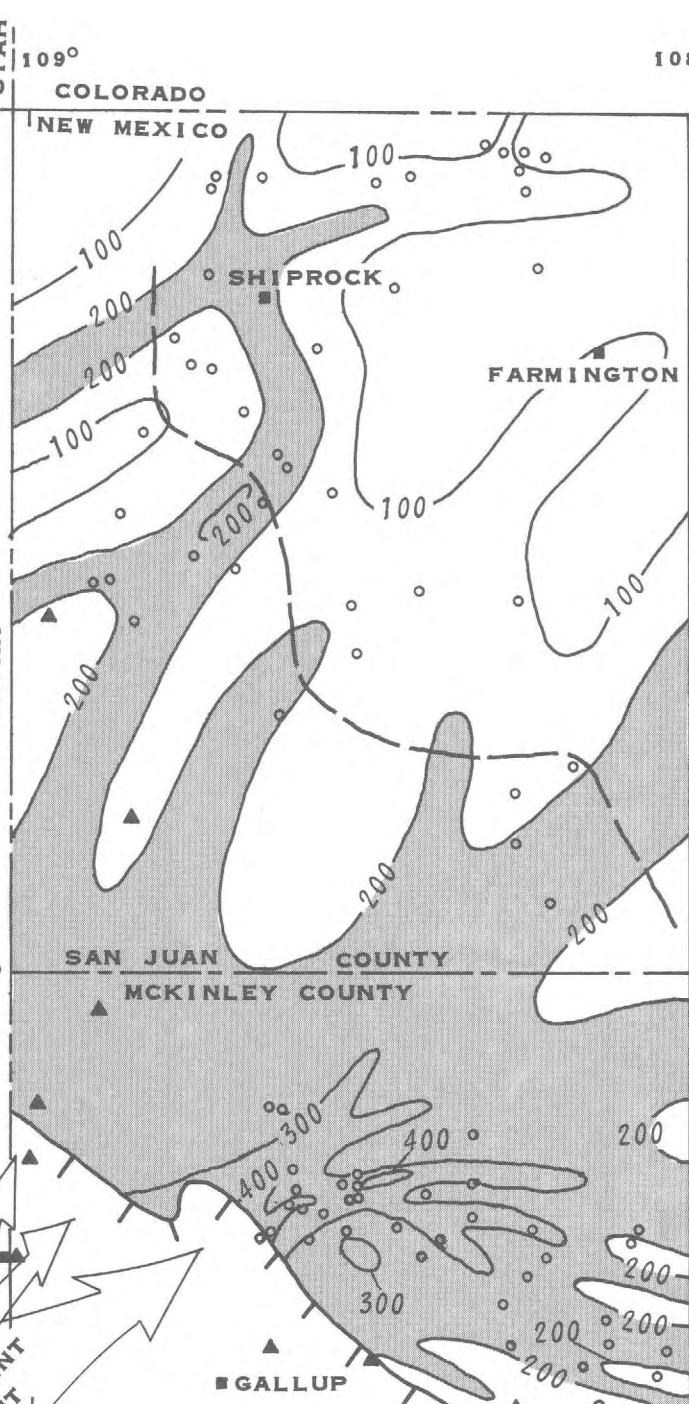


Figure 2.--Generalized aggregate thickness of sandstone in the Westwater Canyon Member of the Morrison Formation based on subsurface well logs and measured outcrop sections (modified from Galloway, 1980, fig. 1).

| Stratigraphic Name | North-West | South-West | South | General Thickness and Lithology of the Morrison Formation Members |
|-------------------------|------------|------------|-------|---|
| | | | | |
| Dakota Sandstone | P | P | P | Lower Basin Member (0-150 feet) Interstratified unit of distinct horizontal sandstone layers, weathering to poorly exposed slopes, deposited in a complex fluvial-lacustrine environment. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Barro Canyon Sandstone | P | P | P | Lower Basin Member (150-400 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Brushy Basin | P | P | P | Lower Basin Member (400-600 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Westwater Canyon Member | P | P | P | Lower Basin Member (600-800 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Recapture Member | P | P | P | Lower Basin Member (800-1000 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Salt Wash Member | P | P | P | Lower Basin Member (1000-1200 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Cow Springs Sandstone | P | P | P | Lower Basin Member (1200-1400 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Horse Mesa Member | P | P | P | Lower Basin Member (1400-1600 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Beckwith Member | P | P | P | Lower Basin Member (1600-1800 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Todito Limestone Member | P | P | P | Lower Basin Member (1800-2000 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Upper sandy member | P | P | P | Lower Basin Member (2000-2200 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Medial silty member | P | P | P | Lower Basin Member (2200-2400 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |
| Lynville Member | P | P | P | Lower Basin Member (2400-2600 feet) Thinbedded to south part of study area (fig. 1, 11a and 11b). Thinbedded sandstone with thin layers of shale and siltstone. Locally conglomeratic, moderately to poorly sorted, friable to finely cemented, thin to medium bedded. Used generally for ground-water development. |

Figure 3.--Stratigraphic nomenclature of the Morrison Formation and adjacent geologic units in the western part of the San Juan Basin (modified from Molenaar, 1977; Condon and Peterson, 1986; and Condon and Huffman, 1988).

GEOLOGY

The Morrison Formation, which crops out along much of the periphery of the San Juan Basin, is overlain by successively younger rocks toward a deep trough in the northeastern part of the basin (Kelley, 1957, fig. 1). During Late Jurassic time, coarser grained sediments (fig. 2) were deposited in the south and west (Santos and Turner-Peterson, 1986, fig. 4). As a result, rocks of Jurassic age generally are 300 to 400 feet thinner in the northeastern part of the San Juan Basin than in the southern and western parts (Santos and Turner-Peterson, 1986, fig. 4). However, accurate thicknesses are difficult to determine in many places where the base of the formation is indistinct on geophysical logs. Geophysical logs do indicate the thickness of the Morrison Formation to range from 500 feet near Gallup to 1,000 feet in the subsurface where there has been no extensive pre-Dakota erosion.

The names and distribution of Jurassic rock units and formations that are stratigraphically below and above the Morrison Formation are shown in figure 3. The Morrison Formation consists of, in ascending order, the Salt Wash, Recapture, Westwater Canyon, and Brushy Basin Members (fig. 3). The Salt Wash Member pinches out between San Jose and Tostitos in the west-central part of the area (fig. 4) (Condon and Peterson, 1986, fig. 4a). The Recapture and Westwater Canyon Members are of greater extent than the Salt Wash and Brushy Basin Members and have been mapped to about 35 miles south of Gallup (Saucier, 1967, figs. 1 and 2). Pre-Dakota erosion has removed the Brushy Basin Member in the southwestern part of the study area (fig. 4).

The Morrison Formation consists of mudstone, claystone, siltstone, sandstone, conglomeratic sandstone, and minor limestone that were deposited in fluvial, eolian, and lacustrine environments. Lithologic descriptions of the four members of the Morrison are shown in figure 3. Colors of the lithologic units in the Morrison Formation are various shades of red, brown, yellow, gray, and green and generally are not characteristic of one rock unit throughout the entire area.

The major structural features in the study area of the San Juan Basin (figs. 1 and 5) are the following: Zuni uplift, Chaco slope, and Gallup sag in the south; Defiance uplift and Defiance monocline in the west; Four Corners platform and Hogback monocline in the central basin; and the Grants area in the east. Several geologic formations in the anticlinal structures on the Four Corners platform contain petroleum. Little or none of the oil and gas produced from these structures, however, comes from the Morrison Formation (Richard Wilson, U.S. Bureau of Land Management, oral commun., 1987).

The general structure of the Morrison Formation is shown by contours drawn within an interval near the top of the Morrison Formation and the base of the Dakota Sandstone in figure 4. The configuration probably is accurate to about 50 to 100 feet. Santos and Turner-Peterson (1986, p. 27-34) and Thaden and Zech (1986, p. 35-46) discussed the problems in mapping the top of the Morrison Formation. The depth below land surface to the top of the Morrison Formation in the Gallup sag near Gallup (fig. 4) is 2,000 feet and near Tohatchi it is 2,700 feet. The depth generally is less than 2,000 feet in the Four Corners platform. In the central basin, the depth to the top of the Morrison Formation near Bisti is 5,000 feet and the maximum depth in the study area (east of Farmington) is 6,700 feet.

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GEOLOGY

The generalized potentiometric surface and distribution of specific conductance of water in the Morrison Formation (fig. 5) were mapped using data collected from 1948 to 1986 (table 1). The long timeframe for data compilation was used because of the lack of sufficient data in a shorter timeframe in much of the study area and because of the lack of stress in most of the area, which would have made the use of the long timeframe inappropriate. Figure 5 is generalized, but it is a useful indicator of hydrologic conditions. The potentiometric contours indicate that the Chuska Mountains and the Zuni uplift are principal areas of ground-water recharge. Additional ground water probably flows into the Morrison Formation in the study area from the northeastern part of the San Juan Basin (Lyford and others, 1980, fig. 2).

Potentiometric contours (fig. 5) indicate that ground water flows out of the study area in the northwestern, southeastern, and probably the southwestern parts of the San Juan Basin. Water is pumped from the Morrison Formation through wells in the Grants uranium region between Navajo Highway 9 and the Zuni uplift in the southern part of the area and from the Four Corners platform between Two Grey Hills and the San Juan River in the northwestern part of the area (fig. 5).

The total pumpage from the Morrison Formation in the study area for municipal, agricultural, and domestic uses and for uranium-mine dewatering has been large (Rearne, 1977, fig. 2; Lyford and others, 1980, fig. 5). Withdrawals from numerous free-flowing artesian wells (Davis and others, 1963; Stone and others, 1983, table 1) and mineral-test holes (Kirk and Condon, 1986, pl. C) have contributed to the lowering of the potentiometric surface in the Morrison Formation.

The practice of completing wells in two or more aquifers (Stone and others, 1983, table 1) to obtain maximum production has created a condition in which water containing large concentrations of dissolved solids might contaminate an aquifer containing water with small concentrations of dissolved solids. The potentiometric-contour closures in the southeast are believed to be the effects of municipal pumping at Crownpoint and Standing Rock and uranium-mine dewatering in the large cone of depression northeast of Church Rock have been rising since the mines in that area closed in 1986.

Most of the water in the Morrison Formation that is available to wells probably is contained in sandstone of fluvial origin that was deposited by braided streams, which is dominant in the Westwater Canyon Member (fig. 5). Although this sandstone, which was deposited by moderate- to high-energy streams (Kirk and Condon, 1986, p. 111), is poorly to well sorted (Turner-Peterson, 1986, p. 32), it can transmit water. The potentiometric contours indicate a primary permeability that is moderate to substantial. This condition is conducive to extensive flow of ground water around and past the thin lenses of siltstone and mudstone (Condon and Peterson, 1986, p. 22) incorporated in the sandstone matrix.

The proportion of fluvial-derived sandstone decreases in the Salt Wash, Recapture, and Brushy Basin Members of the Morrison Formation. On the basis of the proportion of fluvial-derived sandstone, the Westwater Canyon Member has the greatest potential for ground-water development. Condon and Peterson (1986, fig. 5) identified the upper Cow Springs Sandstone as a major eolian facies of the Recapture Member. If fracturing is as extensive in the San Juan Basin as indicated by Kelley (1957, fig. 2), any member of the Morrison Formation could yield water locally.

All except two or three of the wells listed in table 1 probably are completed in the Westwater Canyon Member of the Morrison Formation. The yields of wells completed in the Morrison probably range from a few to about 500 gallons per minute. According to Frenzel and Lyford (1982, fig. 6), transmissivity values of the Morrison Formation range from less than 50 to about 400 feet squared per day in the western part of the San Juan Basin. These values are based on a few aquifer and specific-capacity tests, most of which were concentrated in the northwestern and southeastern parts of the study area. Although data are few, the potential for developing large ground-water supplies from the Morrison Formation appears to be greatest in the large area of thick fluvial-sandstone deposition (fig. 2) between Townships 18 and 24 N. Well 23 near Tohatchi (fig. 5) flowed at the rate of 508 gallons

per minute in 1983, and well 25 near Bisti flowed at the rate of about 400 gallons per minute in 1973. Completion and flow records for the well near Tohatchi indicate that most if not all of the present discharge is from the Westwater Canyon Member. Records for the well near Bisti also identify the Westwater Canyon Member as the source of water.

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