### Figure 1.--Potentiometric surface of the Sparta aquifer, May through June 1989.

## INTRODUCTION

The Sparta aquifer is the principal source of ground water in northcentral Louisiana. In 1985, the aquifer was extensively pumped for public supply (25 Mgal/d) and industrial use (29 Mgal/d, and 7 Mgal/d for other uses) (J.K. Lovelace, U.S. Geological Survey, written commun., 1989). More than 100 public supply systems, in 8 parishes, obtain water from the Sparta aquifer.

Large industrial pumpage from the Sparta aquifer began in 1922 at Bastrop (Sanford, 1973a, p. 60) and in about 1923 at West Monroe. Water levels in wells in the Sparta aquifer have been declining in these areas and in other parts of north-central Louisiana since the early 1920's, when industries began withdrawing large amounts of water. However, in Morehouse Parish the water levels in wells have been recovering since 1982 as a result of a 5 Mgal/d reduction in pumpage at Bastrop (J.K. Lovelace, U.S. Geological Survey, written commun., 1989).

Additional knowledge about ground-water flow and the effects of withdrawals on the Sparta aquifer is needed for assessment of groundwater development potential and protection of the resource. Potentiometric contour maps are used to determine direction of groundwater flow, ground-water gradients, and the effects of pumping on an aquifer system. The rate of ground-water movement can be estimated when the hydraulic conductivity of an aquifer is known.

This report presents data and maps that illustrate the potentiometric surface and water-level changes for the Sparta aquifer and is the third in a series of map reports that show potentiometric surface and water-level changes of aquifers in Louisiana (Fendick, 1989; Fendick and Nyman, 1987). These maps were prepared in cooperation with the Louisiana Department of Transportation and Development. Reports prepared as part of previous studies of northern Louisiana with emphasis on the Sparta aquifer are included in the Selected References.

These maps show the potentiometric surface in the spring of 1989 and water-level changes from 1980 to 1989 for the Sparta aquifer. The regional potentiometric surface of the Sparta aquifer, based on waterlevel measurements made during May through June 1989, is shown in figure 1. Ground-water flow directions are normal to the contours, from higher to lower head.

Water-level changes in wells in the Sparta aquifer from May 1980 to May 1989 are shown in figure 2. Three hydrographs (figs. 3-5) show general water-level trends for periods 1920-89, 1946-89, and 1970-89.

For additional information

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### **GEOHYDROLOGY**

The Sparta aquifer (Sparta Formation) of Eocene age in northern Louisiana consists of very fine to medium sand, silty clay, lignite, and lesser amounts of clay. The Sparta aquifer is comprised of a sequence of alternating sand and clay beds between the massive clays of the overlying Cook Mountain and the underlying Cane River confining units. Sands in the Sparta aquifer were deposited by shifting streams on a deltaicfluvial flood plain (Payne, 1968, p. 3-4). These sands are mostly interconnected, but separately identifiable sands can be traced for only short distances (Snider and others, 1972, p. 13).

The Sparta aquifer generally dips to the east and southeast from the Sabine Uplift. (The Sabine Uplift is shown in figs. 1 and 2.) However, the dip along the northern border of the State is northeasterly. The Sparta aquifer is 50-300 ft thick within the recharge area and thickens to nearly 700 ft near the downdip limit of freshwater. The approximate downdip limit of freshwater is shown in figures 1 and 2. Below the downdip limit of freshwater, all sands in the Sparta aquifer contain saltwater; updip from the freshwater limit, sands in the upper part of the aquifer contain freshwater but sands in the lower part of the aquifer might contain saltwater.

### POTENTIOMETRIC CONTOUR MAP

A regional potentiometric-surface map of the Sparta aquifer in northern Louisiana for May through June 1989 (fig. 1) was constructed using water-level measurements (table 1) made in wells screened in the middle or lower part of the aquifer. The water-level data were adjusted to sea level. The altitude of water levels in wells in the Sparta aquifer ranged from about 160 ft below sea level in an industrial area in West Monroe, Ouachita Parish, to more than 200 ft above sea level in the recharge area in central Bienville Parish. The lowest water levels in the Sparta aquifer are in Ouachita, Lincoln, and Union Parishes.

The natural ground-water flow in the Sparta aquifer in northern Louisiana generally is downdip in an easterly direction from the recharge area in the west toward the Mississippi River alluvial valley (Ryals, 1980). However, the general direction of ground-water movement (fig. 1) indicates that heavy pumpage at West Monroe in Ouachita Parish has altered or reversed the natural direction of flow. Flow in these areas is toward cones of depression in the potentiometric surface at the centers of pumping. Smaller cones of depression occur at pumping centers near Hodge in Jackson Parish, at Ruston in Lincoln Parish, at Winnfield in Winn Parish, and at Minden in Webster Parish. These depressions indicate minor alterations in the direction of flow.

LOUISIANA GROUND-WATER MAP NO. 3: POTENTIOMETRIC SURFACE, 1989, AND WATER-LEVEL CHANGES, 1980-89, OF THE SPARTA AQUIFER IN NORTH-CENTRAL LOUISIANA

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# Table 1.--Water-level data used to construct potentiometric surface of the Sparta aquifer, May through June

	Water level										
Well	Well depth	Water level (feet below	(feet above or below	Date	We:		Well depth	Water level (feet below	(feet above or below	Date	
number	(feet)	land surface)	sea level)	measured	numl	oer	(feet)	land surface)	sea level)	measure	
		Bienv	rille		-			Ouacl	nita		
Bi- 2	490	89.52	170.48	5-11-89	Ou-	24	688	129.32	-44.32	5- 9-8	
Bi- 11	195	17.40	152.60	5- 9-89	Ou-	77	851	236.13	-161.13	5-10-8	
Bi- 13	462	134.93	190.03	5- 9-89	Ou-	401A	397	97.96	-35.68	5-10-8	
Bi- 76	110	49.38	230.62	5-11-89	Ou-	402	750	50.40	12.90	5-10-8	
Bi- 100	90	28.40	191.60	5-11-89	Ou-	404	685	77.75	-16.86	5-10-	
Bi- 112	348	113.45	101.55	5- 9-89		406	681	144.93	-78.27	5-10-8	
Bi- 141	660	242.60	107.40	5- 9-89	Ou-	444	670	214.45	-96.45	5-11-8	
Bi- 144	630	226.62	93.38	5-11-89	Ou-	488	795	319.83	-39.23	5-10-8	
Bi- 166	472	158.32	101.68	5-12-89	-						
Bi- 186	216	39.41	140.59	5- 9-89				Uni	on		
Bi- 192	153	70.42	214.58	5-11-89					16.00	F 10 C	
Bi- 216	300	15.38	184.62	5- 9-89	Un-	2	581	241.02	-16.02	5-10-8	
					Un-	26	745	179.31	-45.39	5- 9-8	
		Cald	well		Un-	78	727	201.72	-26.72	5- 9-8	
					Un-	79	749	146.74	-28.74	5-11-8	
Ca- 51	500	161.33	100.30	5-10-89	Un-	80	710	225.40	-35.40	5-11-8	
Ca- 86B	545	79.70	76.45	5-10-89	Un-	83	326	156.25	-34.75	5-10-8	
					Un-	84	696	233.30	-23.30	5- 9-8	
		Claib	orne		Un-	86	655	92.36	-2.36	5- 9-8	
~. ^		200 55	BC 45	- 10 00°	Un-	134	585	290.99	-70.90	5-11-8	
C1- 9	670	283.55	76.45	5-10-89				Uobo	ton		
C1- 58	488	141.33	108.67	5-10-89				Webs	cer		
Cl- 111	570	280.34	29.66	5-11-89					100.00	F 10 (	
C1- 116	684	246.42	-1.42	5-10-89	Wb-		323	44.91	180.09	5-10-8	
C1- 136	835	316.38	88.62	5-11-89	Wb-		311	52.36	192.64	5-10-8	
Cl- 137	639	224.84	55.16	5-10-89		168	245	62.39	171.61	5-10-8	
C1- 148	625	185.77	4.23	5-10-89		219	136	3.59	186.41	5-11-8	
Cl- 149	736	275.75	-45.75	5-10-89		255	348	102.45	147.55	5-11-8	
C1- 151	719	193.03	-13.03	5-10-89		271	220	82.45	195.55	5-11-8	
		7.1				285	608	179.10	160.90	6-01-8	
		Jack	ison			326	470	113.25	146.75	5-10-8	
		110.15	10.00			338	425	92.71	137.29	5-10-8	
Ja- 49	570	140.17	19.83	5-12-89		341	470	92.10	147.90	5-10-8	
Ja- 88	445	290.33	14.67	5-24-89		349	263	75.29	104.75	5-10-8	
Ja- 135	380	163.60	16.40	6-23-89		357	167	15.18	149.82 200.02	5-10-8 5-11-8	
Ja- 147	703	233.75	-13.75	5-12-89		359	154	79.98	161.32	5-10-8	
Ja- 148	578	214.00	31.00 18.97	5-12-89	wb-	399	298	43.68	101.32	3-10-6	
Ja- 149	478	176.03	10.97	5-12-89	-			Wi	nn		
		Lincoln							71 00	E 10	
					W-	16	583	63.62	71.38	5-12-8	
L- 26	686	168.23	-13.23	5-11-89	W-	76	300	170.60	147.40	5-12-8	
L- 33	811	337.68	-57.68	5-16-89	W-	144B	550	30.72	109.28	5-12-	
L- 68	770	227.34	-47.34	5-12-89	W-	156	829	199.19	57.81	6- 9-8	
L- 113	750	306.36	48.64	5-11-89	W-	164	210	93.72	102.28	5-12-8	
L- 117	551	120. 5	-34.05	5-11-89	W-	172	655	83.88	52.12 131.95	5-12-8 5-12-8	
L- 137	600	281.28	-41.28	5-16-89	W-	177	695	53.05		5-26-8	
L- 143	6 5	271.59	-53.59	5-12-89	W-	178	598	67.70	67.30	5-12-8	
L- 153	535	227.22	52.78	5-10-89	W-	179	585	101.79	93.21	3-12-0	
		Moreh	nouse								
Mo- 5	860	148.14	-30.70	5- 9-89							
Mo- 342	620	89.62	-1.28	5- 9-89							
Mo- 345	786	104.48	9.17	5- 9-89							
Mo- 347	833	162.55	-20.28	5- 9-89							
Mo- 350	740	103.14	8.94	5- 9-89							