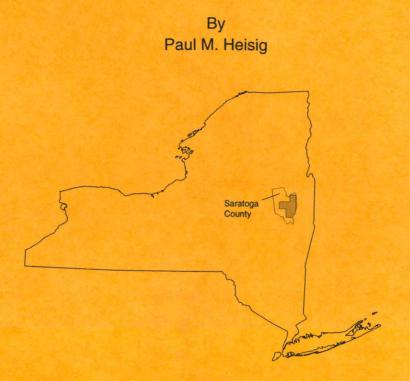
Generalized Stratigraphy, Surficial Geology, Types of Aquifers, and 1988-89 Ground-Water Pumpage in Eastern Saratoga County, New York



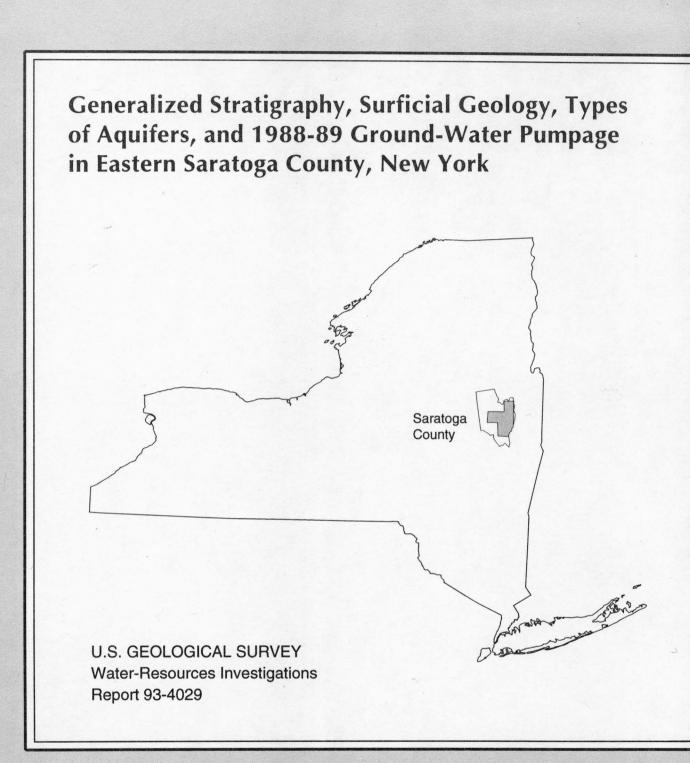
U.S. GEOLOGICAL SURVEY WATER-RESOURCES INVESTIGATIONS REPORT 93-4029

Plates 1-8. Maps showing surficial geology, locations of buried bedrock channels, and locations of ground-water pumping facilities, Saratoga County, N.Y.

- 1. Gansevoort quadrangle
- 2. Quaker Springs quadrangle
- 3. Saratoga Springs quadrangle
- 4. Middle Grove quadrangle
- 5. Round Lake quadrangle
- 6. Mechanicville quadrangle
- 7. Niskayuna quadrangle
- 8. Troy North quadrangle



Prepared in cooperation with the SARATOGA COUNTY ENVIRONMENTAL MANAGEMENT COUNCIL Albany, New York 1994



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GENERALIZED STRATIGRAPHY, SURFICIAL GEOLOGY, TYPES OF AQUIFERS, AND 1988-89 GROUND-WATER PUMPAGE IN EASTERN SARATOGA COUNTY, NEW YORK

By Paul M. Heisig

U.S. GEOLOGICAL SURVEY

Water-Resources Investigations Report 93-4029



Prepared in cooperation with the

SARATOGA COUNTY ENVIRONMENTAL MANAGEMENT COUNCIL

Albany, New York

1994

DEPARTMENT OF THE INTERIOR

BRUCE BABBITT, Secretary

U.S. GEOLOGICAL SURVEY

Gordon P. Eaton, Director

For additional information write to:

U.S. Geological Survey P.O. Box 1669 Albany, NY 12201 Copies of this report may be purchased from:

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Multiply	By	To obtain
	Length	
inch (in)	2.54	centimeter
inch (in)	25.4	millimeter
foot (ft)	0.3048	meter
mile (mi)	1.609	kilometer
	Area	
acre	0.4047	hectare
/	0.09294	square meter
square mile (mi ²)	2.59	square kilometer
	Volume	
gallon (gal)	3.785	liter
	Flow	
foot per second (ft/s)	0.3048	meter per second
foot per day (ft/d)	0.3048	meter per day
cubic foot per second (ft ³ /s)	0.02832	cubic meter per minute
gallon per minute (gal/min)	3.7848	liter per minute
gallon per day (gal/d)	0.003785	cubic meter per day
million gallons per day (Mgal/d)	0.04381	cubic meter per second

CONVERSION FACTORS

Sea level: In this report "sea level" refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929)—a geodetic datum derived from a general adjustment of the firstorder level nets of both the United States and Canada, formerly called Sea level datum of 1929.

GENERALIZED STRATIGRAPHY, SURFICIAL GEOLOGY, TYPES OF AQUIFERS, AND 1988-89 GROUND-WATER PUMPAGE IN EASTERN SARATOGA COUNTY, NEW YORK

By

Paul M. Heisig

ABSTRACT

The eastern part of Saratoga County, N.Y., has undergone considerable population growth and development since 1963, when the completion of Interstate 87 increased accessibility to the Capital District, 30 miles to the south. Since then, nearly all housing developments along the Interstate 87 corridor have obtained their water supply from discontinuous Pleistocene-age sand and gravel aquifers. This report describes the generalized stratigraphy of Pleistocene and Holocene (recent) sediments, as well as two general types of aquifer (bedrock and unconsolidated), and presents data on 1988-89 ground-water pumpage in eastern Saratoga County. Surficial geology, locations of preglacial buried bedrock channels, and ground-water pumpage by facilities that withdraw at least 10,000 gallons per day are indicated on eight 1:24,000-scale topographic $7^{1/2}$ -minute quadrangle sheets (Gansevoort, Quaker Springs, Saratoga Springs, Middle Grove, Round Lake, Mechanicville, Niskayuna, and Troy North). In general, water suppliers obtain more water from unconsolidated aquifers than from bedrock aquifers because the unconsolidated aquifers are more productive. The most productive confined aquifers in the area are kame and lacustrine deltaic deposits, and the most productive unconfined aquifers are outwash or alluvial deposits that are hydraulically connected to surface-water bodies. Lacustrine sand forms a thin, but extensive, aquifer best suited for pumping by shallow wells. The water-resource potential of exposed kame and buried-channel deposits north of Saratoga Lake is largely unknown.

INTRODUCTION

Population growth and development along the Interstate 87 corridor in Saratoga County for the last 30 years has increased the use of ground water for community water supply and, to a lesser extent, for industry. Much of this growth is a result of the completion of Interstate 87 in 1963 (fig. 1), which increased accessibility to the Capital District (cities of Albany, Troy and Schenectady). Although public water supply in Saratoga County is drawn from both surfacewater and ground-water sources, nearly all new development has used ground water as the sole source of supply. Thus, information on groundwater resources in this area is necessary for development and protection of present and future supplies.

In response to the county's need for hydrologic information, the U.S. Geological Survey (USGS), in cooperation with the Saratoga County Environmental Management Council, began a study in 1988 to compile and analyze data on the ground-water resources in eight $7^{1/2}$ minute quadrangles that encompass the eastern part of the county (fig. 2). The investigation entailed the compilation of surficial geology, delineation of preglacial bedrock channels, and compilation of well-depth and pumpage data.

Physical Setting

The study area is in the northern Hudson-Mohawk Lowlands (fig. 1). The area is characterized by low relief and by deposits of Pleistocene age as thick as 300 ft (Reynolds, 1985) that overlie a dissected bedrock surface that was modified by glacial erosion. The deepest of several bedrock channels within the study area is the Colonie channel (Simpson, 1949), which extends to below sea level in places (Dineen and Hanson, 1983; Dineen, 1975-88). Bedrock in the

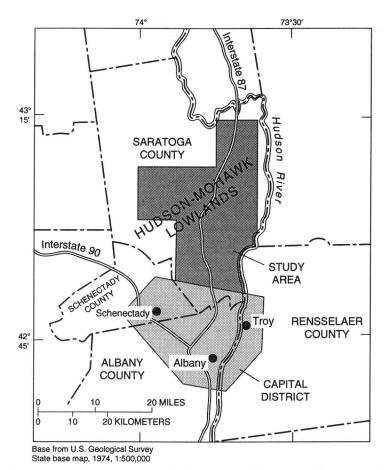


Figure 1.— Major geographic features of Saratoga County and vicinity, and location of study area.

area consists of shale and graywacke of Ordovician age, except in the northern and eastern parts of the study area, where it consists of older sandstone, carbonate, and crystalline rock. Most surface water drains eastward into the Hudson River through three drainage systems: (1) Kayaderosseras Creek, Saratoga Lake, and Fish Creek, (2) Ballston Lake, Ballston Creek, Round Lake, and Anthony Kill, and (3) Snook Kill (fig. 2).

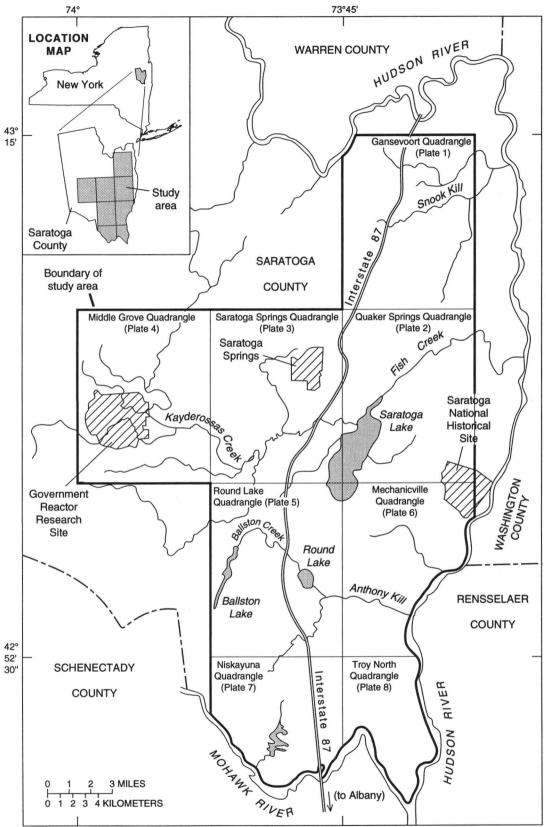
Purpose and Scope

This report discusses the general stratigraphy, surficial geology, types of aquifers, and 1988-89 ground-water pumpage in eastern Saratoga County. The generalized stratigraphic relations of unconsolidated deposits are presented in a block diagram, and well pumpage, well depth, and type of aquifer tapped by yearround suppliers that withdraw at least 10,000 gal/d are tabulated. Plates 1-8 depict the surficial geology, delineate thalwegs (deepest courses) of preglacial buried bedrock channels, and show water-supplier locations and average daily pumpage in the eight 7^{1} /₂-minute quadrangles that form the study area.

Previous Investigations

Glacial geology of the study area has been described by several investigators. Stoller (1916) discussed the glacial geology of the Saratoga quadrangle, and LaFleur (1965, 1979), Dineen and Rogers (1979), and DeSimone and LaFleur (1985) investigated glacial-lake stages, deglacial events, and depositional environments in the Hudson-Mohawk lowlands. Fairchild (1917) discussed postglacial features of the upper Hudson valley, and Hanson (1977) reported on the Late Woodfordian drainage history of the lower Mohawk valley.

Heath (1963) summarized ground-water



Base from U.S. Geological Survey, Heath, 1963, fig. 1-3

Figure 2.—Names and locations of eight 7-11/2-minute quadrangles and geographic features in eastern Saratoga County that are depicted in plates 1 through 8.

conditions in Saratoga County, and Mack and others (1964), Mack (1963), and Heath and Tannenbaum (1963) studied sites at West Milton and the Saratoga National Historical Park, respectively.

Simpson (1949) first identified and mapped buried preglacial bedrock channels, including the Colonie channel, in the Capital District (south of the study area) and recognized the value of the valley-fill aquifers as potential sources of ground water. Dineen and Hanson (1983) mapped and described bedrock topography and glacial deposits of the Colonie channel as far north as Saratoga Lake. Dunn Geoscience Corporation (1981, 1982) and Reynolds (1985) investigated the hydrogeology of the Clifton Park area of Saratoga County, including the Colonie channel aquifers and their recharge areas. Waller (1983) used test-hole data to describe the ground-water potential of buried valley deposits in the Capital District.

Acknowledgments

The author thanks the New York State Geological Survey, which provided maps of surficial geology and bedrock-surface altitude that were used to establish the location of the buried Colonie channel and associated channels north of Saratoga Lake. Walter Bundy of the New York State Department of Health, Bureau of Public Water Supply Protection, provided much of the 1988-89 pumpage data. Many water suppliers provided well and pumpage data.

GENERALIZED STRATIGRAPHY

The stratigraphy of eastern Saratoga County is characterized by a sequence of Precambrianto Ordovician-age bedrock units that are overlain almost everywhere by unconsolidated Pleistocene and Holocene (recent) deposits of variable thickness and composition.

Bedrock

Four general types of bedrock underlie eastern Saratoga County, including from oldest to youngest, crystalline rocks, sandstone, carbonate rocks, and shale (Heath, 1963). The crystalline rocks are a mixture of igneous and metamorphic rock and are of Precambrian age. The Cambrian-age sandstone unit, which contains dolomite interbeds near the top, overlies the crystalline rocks and is 200 to 400 ft thick (Heath, 1963). Carbonate rocks (limestone and dolomite) of Ordovician age overlie the sandstone and are 300 to 400 ft thick in most places (Heath, 1963). The three aforementioned units crop out along the northwestern edge of the study area. The youngest bedrock unit consists of a series of shales of Ordovician age (300 to 1,000 ft thick). This unit forms the bedrock surface throughout most of the study area (Heath, 1963).

The bedrock surface of the study area has been modified by faulting and erosion. Northeast-trending faults along the northwest edge of the study area (Heath, 1963) account for much of the topographic relief and the exposure of older bedrock units in this area.

The bedrock surface is characterized by several preglacial channels that were enlarged and subsequently filled and buried through glacial processes. Delineation of such channels is important because some of the most productive aguifers in the study area are within the channels. The thalweg, or deepest course, of each channel is depicted, along with approximate elevations, in plates 1 through 5 and 7. (No buried channels have been identified in the areas shown in plates 6 and 8.) Channel delineations south of Saratoga Lake are based on investigations by Reynolds (1985) and Dineen and Hanson (1983); those north and west of Saratoga Lake are based on bedrock-surface maps by Dineen (1975-88). The greatest channel-fill thicknesses are found along the Colonie channel (Simpson, 1949), which trends north-south through the study area (pls. 1-3, 5, and 7) and is characterized by a broad channel with a V-shaped inner gorge (Dineen and Hanson, 1983).

Pleistocene Deposits

Pleistocene deposits provide a partial record of the sequence of events and depositional environments during deglaciation of the study area. These events include retreat of the ice front in conjunction with the formation of a time sequence of successively lower stage proglacial lakes in the Hudson River basin as well as major changes in surface-water drainage. A full account of the deglacial history of the study area is beyond the scope of this report; the reader is referred to LaFleur (1979) and Hanson (1977). Pleistocene deposits in the study area represent several types of depositional environments, some of which extended for many square miles, and some of which were localized. A description of each type of deposit and its depositional environment is given in table 1; the generalized stratigraphy and distribution of Pleistocene and Holocene deposits typical of the study area are illustrated in figure 3. Pleistocene deposits can be grouped into four general categories: till, icecontact material, lacustrine material, and outwash.

Till Deposits

Till is an unsorted mixture of rock fragments ranging in size from clay to boulders. It commonly overlies the bedrock surface and is probably the most extensive glacial deposit in the study area. Thickness is generally less than 50 ft (Hanson, 1977) but locally is as much as 250 ft, especially in buried valleys and drumlinoid features (elongated ridges parallel to ice movement that consist of till, and can be bedrock cored). Till generally is present as a surficial deposit at altitudes¹ above 330 to 400 ft, the maximum stage of glacial Lake Albany (DeSimone and LaFleur, 1985), or where erosion has removed overlying kame or glacial-lake sediments.

Ice-Contact Deposits

Ice contact is a general term for stratified, predominantly coarse-grained sediments that were deposited within, against, or atop stagnant or retreating ice masses. These deposits overlie till or bedrock upon which they were deposited during ice meltout or collapse and typically are locally variable with respect to sorting, coarseness, and thickness. Some common types of icecontact deposits include kames, eskers, kame

¹ In this report, altitudes are given in feet above sea level.

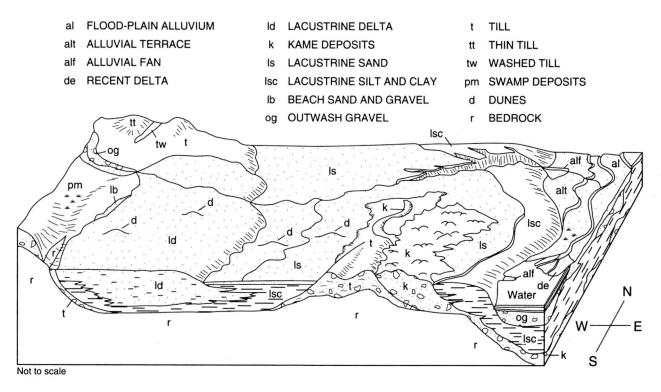


Figure 3.—Generalized stratigraphy of Pleistocene and Holocene (recent) deposits in eastern Saratoga County.

EXPLANATION

terraces, kame deltas, kame moraines, esker deltas, and subaqueous fans. Ice-contact deposits are undifferentiated in plates 1 through 8 to comply with New York State Geological Survey terminology, wherein all ice-contact deposits are identified as "kame deposits." Therefore, the term "kame deposits," as used in this report, refers to undifferentiated ice-contact deposits.

Map Symbol	Description					
	Holocene Deposits					
al	FLOOD-PLAIN ALLUVIUM-recent deposits derived from periodic flooding within a stream valley; generally fine sand and gravel, although sediments commonly reflect the character of the deposits that the stream traverses. Up to 10 ft thick.					
alt	ALLUVIAL TERRACE-glacial and postglacial deposits of fluvial sand and gravel, laterally continu- ous lenses of silt (flood-plain alluvium deposited at a higher level than the recent flood plain). Up to 50 ft thick.					
alf	ALLUVIAL FAN-fan-shaped accumulation of alluvium produced by a stream as it flows from sloped valley walls onto a level flood plain; poorly stratified silt, sand, and cobbles. Up to 20 ft thick.					
de	RECENT DELTA-gravel and sand deposited by tributary streams at a lakeshore. Thickness variable.					
pm	SWAMP DEPOSITS—peat, muck, organic silt, and sand in poorly drained areas. Up to 20 ft thick.					
d	DUNES-fine to medium sand, well sorted, stratified; generally reworked lacustrine sand and deltaic deposits. Up to 50 ft thick.					
	Pleistocene Deposits					
lb	BEACH SAND AND GRAVEL-well-sorted sand and gravel deposited as beaches that were associated with large proglacial lakes.					
ld	LACUSTRINE DELTA-gravel and sand, stratified, generally well sorted, deposited at former glacial lakeshore. Up to 120 ft thick.					
ls	LACUSTRINE SAND-sand deposits associated with large proglacial lakes, formerly a near-shore deposit or subaqueous (underwater) fan. Well sorted, stratified, generally fine to very fine sand. Up to 95 ft thick.					
lsc	LACUSTRINE SILT AND CLAY-Laminated silt and clay deposited in proglacial lakes. Thickness variable but is up to 275 ft in parts of the Colonie channel (Reynolds, 1985).					
fg	FLUVIAL GRAVEL-fluvial sand and gravel, age uncertain; up to 20 ft thick.					
og	OUTWASH SAND AND GRAVEL-sand and gravel deposited by meltwater from retreating glacial ice. Well rounded and stratified, texture becomes increasingly fine away from ice border. Up to 25 ft thick.					
k	KAME DEPOSITS-a stratified mix of cobbles, gravel, sand, silt and clay that was deposited by glacial meltwater in contact with stagnant glacial ice. Large variability in sorting, particle size, and bed thickness. These ice-contact deposits include kames, eskers, kame terraces, kame deltas, esker deltas, and subaqueous fans. Up to 100 ft thick.					
cd	COMPLEX DRIFT-large variability in composition over short distances. Ranges from ice-contact sand and gravel to lacustrine sand, silt, and clay, and till. Forms scarps on either side of the Hudson River flood plain.					
t	TILL-lodgment or ablation till of variable thickness and texture. Up to 250 ft thick. Texture ranges from clayey silt to boulder-bearing clay.					
tt	THIN TILL-variable mantle of rock debris and till of variable thickness. Bedrock sporadically crops out. Generally less than 10 ft thick.					
tw	WASHED TILL—veneers of cobbles and boulders in surface channels and terraces in thick till deposits. Can show imbricate bedding structure. Generally less than 10 ft thick.					

Lacustrine Deposits

 $\langle \hat{e} \rangle$

As the glacial-ice margin retreated northward, meltwater became impounded and formed glacial Lake Albany, which stood at elevations between 330 and 400 ft (LaFleur, 1965; Hanson, 1977; DeSimone and LaFleur, 1985). These elevations represent lake stage near Albany and Moreau State Park (Gansevoort quadrangle), respectively, and indicate a tilted glacial-lake water plane at present, which was caused by rebound of the land surface to the north after retreat of the ice front. The maximum stage (glacial Lake Albany) was followed by several successively lower lake stages (LaFleur, 1965; Hanson, 1977; DeSimone and LaFleur, 1985). The Pleistocene sediments associated with these lake stages include lacustrine silt and clay, lacustrine sand, lacustrine deltas, and lacustrine beach deposits. Silt and clay forms the most extensive and thickest lacustrine deposit in the study area, and its thickness ranges up to 275 ft (Reynolds, 1985). Silt and clay commonly overlies till and kame deposits and, in turn, is typically overlain by lacustrine sand or alluvial deposits. Silt and clay are exposed at land surface where lacustrine sand and alluvial deposits either were not deposited or have been eroded. Most lacustrine sand deposits in this area are 20 to 50 ft thick and have typically been reworked into sand dunes at land surface. In some areas, the sand directly overlies older

Pleistocene ice-contact sediments. Lacustrine deltaic deposits are commonly elevated above, and interfinger with, adjacent lake-bottom sediments. Another lacustrine delta is found above the highest level of glacial Lake Albany, where it was deposited in an impounded water body along the ice margin during deglaciation.

Outwash Deposits

Outwash is a stratified deposit composed primarily of sand and gravel derived from meltwater outflow from a glacier. It is deposited beyond the active margin or the end moraine of the glacier. Outwash is present as a surficial deposit along the Kayaderosseras Creek (Middle Grove quadrangle, pl. 4), which is higher in elevation than the maximum stage of glacial Lake Albany.

Holocene (recent) Deposits

Recent deposits in the study area consist of alluvium, alluvial-terrace, alluvial-fan, dune, and swamp deposits. Alluvial deposits most commonly overlie the Pleistocene or glacial deposits and form flood plains along stream or river courses. In some areas, extensive lacustrine delta and lacustrine sand deposits became exposed after Lake Albany drained to lower levels and then were reworked by wind into dunes and dune fields.

SURFICIAL GEOLOGY

The surficial geology of the study area is depicted on plates 1 through 8. The distributions of surficial units form generalized patterns that are consistent on all of the maps; namely that till and bedrock are most commonly exposed in the uplands, kame-deposit exposures generally border the uplands, and lacustrine-deposit exposures predominate in areas of low elevation and low relief. The boundary between lacustrine deposits and other surficial material is typically at elevations of 330 to 400 ft, the highest stage of glacial Lake Albany, but isolated lacustrine deposits can be found at higher elevations wherever meltwater drainage was locally impounded by ice or glacial drift. Till, kame deposits, and bedrock are present at land surface locally at low elevations wherever the overlying lacustrine deposits have been eroded. An example of a lacustrine deposit at a high (above 400 ft) elevation is the Milton delta (Stoller, 1916), a deltaic deposit that lies at 420 ft within the Saratoga Springs and Middle Grove quadrangles (pls. 3 and 4). An example of till, bedrock, and kame deposits exposed at a low elevation (below 330 ft) is the postglacial valley now occupied by Ballston Lake (pl. 5).

Holocene alluvium is found along presentday drainage courses. Alluvial-terrace deposits are found alongside and, especially, near the mouth of former postglacial drainage courses.

Gansevoort Quadrangle

Except in its northwestern corner and occasional drumlinoid (till) features, lacustrine deposits are predominant in the Gansevoort quadrangle (pl. 1). Sand plains and dune fields derived from lacustrine sand cover the surface and form a broad plain that has been partly dissected by the Snook Kill and its tributaries (fig. 2). Exposures in the dissected areas reveal that the lacustrine sand forms a veneer over a thicker sequence of lacustrine silt and clay. In the northwestern corner of the quadrangle is a bedrock upland (the Palmerton Range), which is separated from the lacustrine deposits by narrow, discontinuous till bodies and a large kame deposit.

Quaker Springs Quadrangle

The Quaker Springs quadrangle (pl. 2) is characterized by two general areas, one of low relief and one of moderate relief, that flank the Saratoga Lake - Fish Creek drainage (fig. 2). The low-relief area that occupies the northwestern part of the quadrangle is predominantly lacustrine sand, although an area just north of Saratoga Lake contains extensive alluvialterrace deposits. Lacustrine silt and clay is not extensively exposed at land surface because the overlying sand plain in this part of the quadrangle is relatively undissected by streams. Southeast of the Saratoga Lake - Fish Creek drainage, deposits of lacustrine and dune sand decrease in number, and till, bedrock outcrops, and lacustrine silt and clay become increasingly common with increasing elevation and relief.

Saratoga Springs Quadrangle

The Saratoga Springs quadrangle (pl. 3) is characterized by a wide variety of surficial deposits. The northwest and north-central parts are at higher elevations and have greater relief than the rest of the quadrangle; thus, till and bedrock outcrops predominate here. A discontinuous band of kame deposits follows the southern boundary of the till and bedrock uplands north and west of Saratoga Springs. Most of the surficial deposits in the rest of the quadrangle consist of lacustrine and alluvial deposits. The lacustrine deposits include sand and gravel, sand, and silt and clay. A lacustrine deltaic deposit, the Milton delta (Stoller, 1916), is the dominant topographic feature in the southwestern quarter of the quadrangle. Lacustrine sand forms a broad plain over the central and northeastern parts of the quadrangle and also caps the Malta Ridge in the southeastern

corner. Lacustrine silt and clay is relatively sparse in this quadrangle and does not underlie the lacustrine sand everywhere.

The area in the southeastern part of the quadrangle has largely been dissected by glacial meltwaters in the area now occupied by Kayaderosseras Creek. Alluvial-terrace and washed-till deposits are common in this area.

Middle Grove Quadrangle

The Middle Grove quadrangle (pl. 4) lies entirely above the elevation of the maximum stage of glacial Lake Albany (330 to 400 ft) and, therefore, lacks extensive deposits of lacustrine sand, silt, and clay. Instead, it is overlain largely by till and kame deposits; till mantles most topographic highs, whereas kame deposits flank the hills and cover much of the land at lower elevations. Two other surficial materials are present within the quadrangle: outwash gravel occupies the Kayaderosseras Creek valley and grades into lacustrine delta deposits of the Milton delta (which formed in a local ice-marginal lake), and many bedrock exposures are found in the northwestern corner of the quadrangle, where elevations exceed 1,000 ft.

Round Lake Quadrangle

Lacustrine sand plains and dune fields cover much of the eastern half of this quadrangle (pl. 5), and till covers the high elevations in the western half. The most prominent features are several postglacial drainage channels that have been incised into the surficial material without subsequent infilling (Dineen and Hanson, 1983). Several present-day drainage courses (Ballston Lake, Ballston Creek, Mourning Kill, Drummond Creek, Round Lake, and Anthony Kill) occupy these postglacial drainage channels. Downcutting of the streambed, followed in some areas by minor postglacial deposition, has resulted in a pattern of surficial deposits parallel to, and exposed along, the channel walls or in the thalweg of the channel. The surficial geology along these channels reflects the local topography: the Ballston Lake, Ballston Creek, and Mourning Kill are characterized by bedrock exposures along the channel walls, whereas Drummond Creek and Round Lake-Anthony Kill channel sections are characterized by till and lacustrine silt and clay.

Mechanicville Quadrangle

The area shown in the southern half of this quadrangle (pl. 6) is covered mostly by lacustrine sand, silt, and clay deposits, which are punctuated by occasional till or bedrock hills or kame deposits. Within this half of the quadrangle, the Anthony Kill valley, a postglacial stream channel, contains till or bedrock outcrops along the valley walls and thin fluvial gravel deposits along its southern edge. The only substantial kame deposits in the quadrangle are north of the Anthony Kill at Willow Glen and in the area east of the Mechanicville Reservoir.

A large alluvial terrace occupies the western side of the Hudson River valley from Stillwater Junction to Bemis Heights. This terrace borders a nearly ${}^{3/}_{4}$ -mi-wide alluvial flood plain just north of Stillwater. The northern part of the quadrangle, west from the Hudson River alluvial deposits, is covered mostly by lacustrine deposits of glacial Lake Albany. The area southwest of Schuyler Creek is overlain mostly by lacustrine silt and clay, and the area northeast of Schuyler Creek contains about equal areas of (1) lacustrine silt and clay and (2) till and bedrock. The northwestern part of this quadrangle contains areas of lacustrine sand that have been reworked to form extensive dune fields.

Niskayuna Quadrangle

The part of the Niskayuna quadrangle (pl. 7) that lies within Saratoga County is dominated by lacustrine sediments associated with glacial Lake Albany. The western half of the mapped area contains till and lacustrine silt and clay and is bordered on the east by a north-south-trending band of lacustrine beach and deltaic deposits. The eastern half of the quadrangle contains extensive lacustrine-sand plains and dune fields. These deposits are underlain in places by older ice-contact sand and gravel deposits that are locally exposed at land surface. The southern limit of the study area (the Mohawk River) is bordered by alluvial-terrace deposits, till, lacustrine silt and clay, and bedrock.

Troy North Quadrangle

Most of the mapped area in this quadrangle (pl. 8) was inundated by glacial Lake Albany, as evidenced by the lacustrine sand, silt, and clay that cover nearly all of the mapped area. Exceptions include sporadic till hills, occasional bedrock outcrops along the Mohawk River, and continuous alluvium along the Hudson River. Alluvium along the west bank of the Hudson River. River forms a flood plain as much as 3/4 mi wide that contains several alluvial fans.

TYPES OF AQUIFERS

Ground water is pumped from unconsolidated deposits and bedrock within the study area; the unconsolidated aquifers are generally more productive and, thus, are more widely used than bedrock aquifers. The hydraulic characteristics of unconsolidated deposits are summarized in table 2. The most productive aquifers are in unconsolidated deposits of outwash, kame, lacustrine delta, and, to a lesser extent, lacustrine sand of Pleistocene age.

Kame Deposits

Kame deposits in preglacial bedrock channels, such as the Colonie channel (Simpson, 1949), are overlain by lacustrine silt and clay and form discontinuous, but productive, confined aquifers ranging up to 100 ft thick (Dineen and Hanson, 1983; Waller, 1983; Reynolds, 1985; Dunn Geoscience, 1981). One such confined aquifer system, the Colonie channel aquifer, supplies water to several developments in the Town of Clifton Park (Dunn Geoscience, 1981; Reynolds, 1985). Well yields as high as 800 gal/min have been reported (Dunn Geoscience, 1981), although most supply wells yield far less as a result of well design, small aquifer extent, pumping limitations imposed by the design of the distribution system, or interference from nearby production wells. Outcrops of kame deposits adjacent to the buried channel probably serve as recharge areas for the confined aquifer (Reynolds, 1985). Confined aquifers are susceptible to contamination by human activities within these recharge areas.

The locations and extent of confined kame deposits are not discernible on plates 1 through 8, but the buried-channel locations provide a starting point for future subsurface investigations. Neither the composition of buried valley deposits north of Saratoga Lake nor the waterresources potential of most surficial kame deposits is known.

Lacustrine Deltaic Deposits

Lacustrine deltaic deposits form productive aquifers in at least two locations in the study area—(1) the Milton delta (labeled ld) west of Saratoga Springs (pl. 3) and (2) a confined deltaic deposit northwest of West Milton (pl. 4). The Milton delta supplies several housing tracts and parts of the City of Saratoga Springs and the Village of Ballston Spa. The largest withdrawals are from the Geyser Crest well field (Saratoga Springs quadrangle, pl. 3), where six wells are screened within 60 ft of land surface and where the maximum individual well yield is reported to be between 600 and 700 gal/min (Dunn Geoscience, 1988). A confined aquifer composed of lacustrine deltaic sand is the principal source of water for the Kenneth A. Kesselring site (fig. 2), a Federal Government reactor-research installation, at West Milton. Studies of this aquifer suggest that yields of 800 gal/min can be obtained from a single well (Mack, 1963).

Lacustrine Sand, Outwash Sand and Gravel, and Alluvium

These deposits typically form thin, unconfined aquifers, and their proximity to land surface makes them more susceptible than confined aquifers to contamination from human activities, such as septic waste and road salt. Lacustrine sand deposits form the most widespread aquifers in the study area. Hydraulic properties of the aquifer at the Saratoga National Historical Park (pls. 2, 6; fig. 2) were investigated by Heath and Tannenbaum (1963),

Table 2.—Hydraulic characteristics of unconsolidated deposits in the study area.

Map Symbol	Description
al	FLOOD-PLAIN ALLUVIUM-Associated with large streams and rivers. Has potential for water supply by induced infiltration of surface water into the aquifer by pumping. Permeabil- ity varies with local sediment character of basin. Alluvium is generally not considered a major source of ground water within small stream valleys, where its thickness and extent is typically limited.
alt	ALLUVIAL TERRACE-Generally permeable; commonly unsaturated.
alf	ALLUVIAL FAN-Generally permeable but of limited extent; commonly unsaturated.
de	RECENT DELTA-Generally permeable but of limited extent.
d	DUNES-Permeable; commonly unsaturated.
\mathbf{pm}	SWAMP DEPOSITS-Generally impermeable.
lb	BEACH SAND AND GRAVEL-Permeable but limited in areal extent and thickness; generally unsaturated.
ld	LACUSTRINE DELTA-Generally permeable and productive where saturated.
ls	LACUSTRINE SAND-Permeable; saturated thickness could limit invidiual well yield.
lsc	LACUSTRINE SILT AND CLAY-Typically impermeable; generally a confining unit where underlain by permeable deposits.
fg	FLUVIAL GRAVEL-Generally permeable; commonly unsaturated.
og	OUTWASH SAND AND GRAVEL-Permeable, with potential for induced infiltration of surface water where in hydraulic connection with streams or rivers.
k	KAME DEPOSITS-Permeable and productive where saturated.
cd	COMPLEX DRIFT-highly variable permeability.
t	TILL-Low permeability; can yield small amounts of water to large-diameter (2 ft or greater) dug wells; water is derived primarily from local sand lenses within the till.
tt	THIN TILL-Low permeability, commonly unsaturated, yields little water.
tw	WASHED TILL-Small areal extent and thickness; can be permeable, depending on the degree of sorting.

who used aquifer-test data to estimate a horizontal hydraulic conductivity of 94 ft/d and a storage coefficient of 0.16. The saturated thickness of these relatively thin aguifers could limit the long-term yield of wells. Several municipal and private subdivision suppliers tap these aguifers; the maximum reported withdrawal from these unconsolidated aquifers within the study area (table 3, pumping site 13; pl. 3) is 250,000 gal/d. Saturated outwash sand and gravel and alluvium in the valley of Kayaderosseras Creek (pl. 4) also represent a potential water source. Although these deposits are less extensive than those previously mentioned, downward gradients caused by pumping can induce flow from the creek and tributary streams into the shallow aquifer and thereby increase the amount of water available for pumping. Water pumped from a shallow collector well near Kayaderosseras Creek at West Milton (pl. 4) supplied the Kenneth A. Kesselring site (fig. 2) for 8 years and was capable of yielding 750 gal/min for extended periods (Mack, 1963), although the maximum yield fluctuated seasonally with the level of Kayaderosseras Creek.

Bedrock

The most productive bedrock aquifers in the study area consist of carbonates and sandstone that form the bedrock surface along the northwestern edge of the study area. Much of the remaining study area is underlain by shale with the most productive zone, in terms of yield and water quality, being a fractured and weathered layer near the bedrock surface. Well yield tends to decrease, and water mineralization tends to increase, with depth. Fractured crystalline rock is used as a source of water in some areas near the northwestern boundary of the study area. Most bedrock aguifers within the study area are confined by till or fine-grained lacustrine sediments. Heath (1963) lists average yields for wells completed in bedrock in Saratoga County as follows: shale, 10 gal/min; carbonate rock, 30 gal/min; sandstone, 20 gal/min; and crystalline rock, 5 gal/min. Well yields from the bedrock, which depend on secondary permeability (fractures), vary widely and are difficult to predict before test drilling and pumping.

GROUND-WATER PUMPAGE IN 1988-89

Pumpage data from suppliers that average at least 10,000 gal/d were compiled from the New York State Department of Health data base and from telephone interviews with suppliers (table 3). The resulting data set includes water supplies for municipalities and private subdivisions, apartment complexes and condominiums, mobile-home parks, private industry, and government facilities; seasonal and commercial pumpage was not compiled. Table 3 lists the supplier; the $7^{1/2}$ -minute quadrangle in which it is located; the average daily pumpage; general aquifer type; and the number, type, and depth of the production wells (where available). The location of each supplier is shown on plates 1 through 8.

Ground-water pumpage for domestic use far exceeds that for industrial use within the study area (table 4). Because many industrial plants require large amounts of water for certain processes, such as cooling, they are commonly built near surface-water bodies (Hudson and Mohawk Rivers, Kayaderosseras Creek), where large quantities of water are readily available.

Of the 48 ground-water suppliers listed in table 3, 69 percent tap unconsolidated aquifers, 21 percent tap bedrock aquifers, and 10 percent tap both types of aquifer. Unconsolidated aquifers are the more productive source, however, and are the source for all withdrawals of more than 100,000 gal/d in the study area.

Water in unconsolidated aquifers can be divided into two general depth categories shallow (less than 50 ft) and deep (greater than or equal to 50 ft). Most withdrawals of shallow ground water are from extensive lacustrine sand deposits, but alluvium, fluvial gravels, and outwash gravels are important shallow sources locally. Other permeable surficial deposits (listed in table 2) can also serve as aquifers, provided that they have sufficient saturated thickness to support the intended pumpage. Most withdrawals of shallow ground water are from unconfined aquifers, and most withdrawals of deep ground water are from confined aquifers in buried valleys or thick deltaic deposits.

Table 3.—Average daily pumpage (1988-89), aquifer type, and well data from suppliers that withdraw at least 10,000 gallons per day, in eastern Saratoga County, arranged by facility,.

[Data from New York State Department of Health and telephone interviews; gal/d = gallons per day. NA Data not available. * Estimated from available data. ** Design capacity upon completion.]

Designatio of pumping site (see pls 1-8)	5	Quad- rangle ^a	Average daily pumpage (gal/d)	General type of Aquifer ^b	Number and types of wells ^c , and well depths, in feet below land surface
MUNICIP	ALITY OR PRIVATE SUBDIVISIO	N			
1	Village of Ballston Spa	SAR	576,000	U	1 well - 80
2	Calico Colony-Woodland Hills	NIS	80,000	U	3 wells - 62, 62, 66
3	Clifton Gardens Water Supply	NIS	150,000	U	15 driven wells - 20
					13 driven wells - 20
	a a				3 wells - 156, 164, 153
4	Colonial Green Water Supply	NIS	50,000	U	3 wells - 53, 83, 80
5	Colonial Hills Water Co.	SAR	98,000 **		2 wells - 205, 205
6	Country Knolls Water Works, Inc.		1,491,000	U	9 sites:
A	Elnora well field	\mathbf{RDL}			12 driven wells-26 to 32
В	Kinns Road well field				1 well-112
C	Pierce Farm well field	RDL			1 well-130
D	Moe Road well field	RDL			1 well-165
E	Lapp Road, well field	NIS			2 wells-132, 132
F	Plank Road, well field	NIS			1 well-185
G	Golf course, well field	RDL			2 wells-90, 90
H	Northcrest well field	NIS			infiltration well-20 to 25
Ι	Sherwood Forest well field	NIS			3 wells -25, 25, 25
	Constant Prototon	RDL	00F 000	TT	4 malla 149 150 194 194
7	Crescent Estates	NIS	325,000	U U	4 wells - 148, 150, 134, 134 6 wells - all between 52 to 57
8	Geyser Crest-Heritage Knolls	SAR	1,013,000	-	
9	Gilbert Road Water Works	QSP	10,000	U	
10	Heritage Springs Water Co.	SAR	30,000	B U	3 wells - 170, 170, 210 3 wells - 155 each
11	Interlaken Water Works	QSP	40,000	U	2 wells - 38 each
12	Laurel Acres	SAR RDL	12,000	U	5 wells - all between 32 to 35
13 14	Luther Forest Park Lane Estates Water Works	RDL NIS	250,000 30,000	U	4 wells - 50 each
14 15	Rowlands Hollow Water Co.	SAR	45,000	U, B	3 wells - 112 and 100 to 120 in uncon
19	Rowlands Hollow water Co.	SAR	40,000	О, Б	idated material, 183 in bedrock
16	Pine North	SAR	12,000	U	1 spring-fed cistern
10	Village of Stillwater	MEC	320,000	Ŭ	5 wells - all between 24 to 39
18	Terel Hills	GAN	30,000	U	3 wells - 87, 92, 90
			50,000		5 weils - 57, 52, 55
PARIME	ENT COMPLEX				
19	Brookline Village	RDL	25,000	В	8 wells - 75 to 150
20	Clifton Park	NIS	24,000	U	3 wells - 70 each
21	English Village	GAN	16,150	U	2 wells - 180 each
22	London Square	NIS	12,000	U	2 wells - 120, 90
23	McGregor Mini-mart	GAN	10,000	U	4 driven wells - 25 to 35
24	Parkwood Village	TYN	14,000	B *	3 wells - 150 each
25	Squire Park	NIS	10,000		3 wells - NA
26	Twin Lakes	NIS	75,000	U	1 well - 81

^a Quadran	gle abbreviations:		
GAN	(Gansevoort)	RDL	(Round Lake)
QSP	(Quaker Springs)	MEC	(Mechanicville)
SAR	(Saratoga Springs)	NIS	(Niskayuna)
MGR	(Middle Grove)	TYN	(Troy North)

^c Well types include:

wells - drilled unless specified otherwise driven wells - well points, typically less than 35 ft deep dug wells - typically less than 25 ft deep infiltration gallery - large-diameter casings, typically less than 25 ft deep

^b U unconsolidated; B bedrock

Table 3.—Average daily pumpage (1988-89), aquifer type, and well data from suppliers that withdraw at least10,000 gallons per day, arranged by facility, in eastern Saratoga County (continued).

Designation of pumping site (see pls 1-8)		Quad- rangle ^a	Average daily pumpage (gal/d)	General type of Aquifer ^b	Nun dept	nber and types of wells ^C , and well hs, in feet below land surface
MOBILE H	OME PARK					
27	Adirondack Estates	GAN	11,000	U	10 w	rells - NA
28	Brookview	SAR	35,000	В	1	well - 100
29	Cobblestone	QSP	22,000	U	3	driven wells - NA
30	Creek and Pines	MGR	35,000	U	2	driven wells - NA
31	D and R	NIS	35,000	U, B	4 8	wells - 2 in bedrock - 70, 235 2 in unconsolidated material - 40 each driven wells 15 to 25
32	Ebert's Trailer Town	SAR	14,000	U, B	3	wells - 97 (probably in bedrock) 74, 79 (probably in unconsolidated deposits)
33	Hidden Acres	SAR	14,000	B	1	
34	Loughberry	SAR	14,300 *	U	4	driven wells - 35 to 45
35	Malta Gardens	RDL	40,000	U	4	wells - 30 each
36	Malta Mobile Acres	RDL	16,000	U	2	dug wells - NA d
37	Martindale Court	NIS	15.000	U	1	well - 35
38	Northern Pines	MGR	40.000	Ū	3	infiltration wells - 15, 15, 25
39	Northway Mobile Estates	RDL	11,000	Ū	6	wells - 100 to 150
40	Pyramid Pines	QSP	30,000	U	13	driven wells - NA
41	Stockade	MGR	20,000	B *	4	wells - 180, 180, 280, 280
42	Terrace Hill Commons	MGR	22,000	B	7	wells - 5 between 240 to 385, 2 are 7
43	Turf	NIS	119,000	U, B	6	wells - 78, 85, four wells between 10 and 170; 2 deepest wells in bedrock
44	Utter's Mobile Estates	SAR	11,000	В	2	wells - 119,121
45	Vosburg Road	RDL	15,200	U, B	2	wells - 115, 160 in bedrock 1 infiltration gallery - 10
46	Wilton Mobile Park	QSP	30,000	U	6	driven wells - between 19-21
PRIVATE I	NDUSTRY					
47	General Foods	SAR	200,000 *	U	1	well - 83
GOVERNM	ENT INSTALLATION					· · · · · · · · · · · · · · · · · · ·
48	Kenneth A. Kesselring Site	MGR	2,600,000	U	7	wells - 105, 99, 100, 72, 25, 25, 122

Type of Supplier	Pumpage,in gallons per day
Municipalities and subdivisions	4,562,000
Apartment complexes and condominiums	186,150
Mobile home parks	549,500
Private industry	200,000
Government installations	2,600,000

Table 4.—Pumpage totals, by supplier (See table 3.)

Every category of supplier listed in table 3 has wells generally less than 50 ft deep that are screened in shallow aquifers, but small suppliers withdraw a larger percentage of their water from shallow aquifers than the large suppliers do. For example, 65 percent of mobile-home parks have wells screened in shallow aquifers, whereas only 33 percent of municipalities and private subdivisions have these wells. The difference in usage reflects differences in:

- 1. Water needs.—Thin, near-surface aquifers meet the needs of small suppliers more readily than those of large suppliers.
- 2. *Economics.*—Shallow wells, particularly driven wells, are less expensive to install than deep, typically large-diameter wells.
- 3. Location.—Several municipalities and private subdivisions are directly above or near the Colonie buried bedrock channel, whereas most mobile-home parks are scattered about the study area, away from known confined aquifers in these channels.

Most of the deep wells that are screened in unconsolidated aquifers belong to municipal and private subdivision suppliers. The most common sources of water at these depths are saturated kame and lacustrine deltaic deposits, and most production wells screened in these deposits are between 100 and 200 ft deep.

SUMMARY

Aquifers are the sole source of water for most domestic development in eastern Saratoga County; therefore, information on the stratigraphy, surficial geology, types of aquifers, and groundwater pumpage can be helpful for development and protection of this resource. The USGS, in cooperation with the Saratoga County Environmental Management Council, began a study in 1988 to compile and analyze data on the groundwater resources of the eastern part of the county.

The study area is underlain by a dissected bedrock surface that has been further modified by erosional processes associated with glaciation. Pleistocene and Holocene (recent) deposits blanket much of the bedrock surface and fill the preglacial drainage channels. The glacial deposits can be divided into the general categories of till, ice-contact, lacustrine, and outwash deposits. Lacustrine deposits are the predominant surficial material at elevations below the maximum glacial-lake stage, which increases in elevation to the north; till and ice-contact deposits predominate above that level but also are found beneath the lacustrine deposits. The study area contains aquifers that consist either of unconsolidated sediments or bedrock. The unconsolidated sediments, such as kame and lacustrine delta deposits, form the most productive aquifers, which can be confined or unconfined. Lacustrine sand and outwash deposits form productive unconfined aquifers near land surface, but they are more susceptible to contamination from surface sources than confined aquifers. In general, bedrock aquifers are less productive than the unconsolidated aquifers, are typically confined, and contain more highly mineralized water.

Most of the pumped ground water is from unconsolidated, Pleistocene-age aquifers. The primary use is for domestic purposes; private industry uses mostly surface-water sources. Lacustrine delta and buried-channel kame deposits are tapped by the largest ground-water users; aquifers composed of lacustrine sand and, to a lesser extent, aquifers composed of outwash, are pumped in parts of the study area where the more productive aquifers are absent or where water demands are low.

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HEISIG Generalized stratigraphy, surficial geology, types of aqiufers, and 1988-89 ground-water pumpage in Eastern Saratoga County, New York USGS/WRI 93-4029