

#### ABSTRACT

Water-level measurements were made during 1992-94 in 253 wells completed in the Black Creek aquifer in an approximately 2,550-square-mile area of Bladen, Hoke, Robeson, and Scotland Counties in the southern Coastal Plain of North Carolina. Water levels were measured in 56 wells in the fall of 1992, 135 wells in the fall of 1993, and 62 wells in the fall of 1994 to map the potentiometric surface of the Black Creek aquifer for these different parts of these counties. The maps of the potentiometric surface can be used to infer general direction of ground-water flow from recharge areas in the uplands to discharge areas at local streams and wetlands. Withdrawals from wells at pumping centers, such as areas around Elizabethtown in Bladen County and Maxton and Lumberton in Robeson County, have resulted in cones of depression in the potentiometric surface, causing ground-water to flow toward the pumped wells. In southwestern Bladen County, where the Black Creek aquifer is overlain by the Peedee aquifer and the Black Creek confining unit, most ground water does not discharge to local streams but flows southeastward toward the coast or is discharged from pumped wells.

Observed water-level changes in the Black Creek aquifer have been minor throughout most of Bladen, Hoke, Robeson, and Scotland Counties since 1988. However, water-level declines of about 11 feet were recorded near Lumberton during 1993-94.

#### INTRODUCTION

In 1992, the U.S. Geological Survey (USGS) began a 3-year investigation of water levels in the major water-supply aquifers in the northern Coastal Plain of North Carolina. The study was conducted in cooperation with the Lumber River Council of Governments (LRCOG), which comprises local governments from Bladen, Hoke, Robeson, and Scotland Counties (fig. 1). The objective of the investigation was to assess the effects of ground-water withdrawals on water levels in major aquifers in the LRCOG area, an area covering about 2,550 square miles.

This report describes water-level conditions in the Black Creek aquifer in parts of Bladen, Hoke, Robeson, and Scotland Counties, North Carolina. The results include an assessment of the effects of ground-water withdrawals on the potentiometric surface of the Black Creek aquifer during the fall of 1992, the fall of 1993, and the spring of 1994. Selected hydrographs illustrate seasonal variations and the effects of ground-water withdrawals.

The approach for the investigation was to divide the counties in the study area of the LRCOG into three study units (fig. 2). A different unit was selected for study during each year of the 3-year investigation. Water-level data for the first study unit were collected in parts of Robeson and Bladen Counties during the fall of 1992 and reported by Strickland (1994). Water-level data for the second study unit, which included parts of Bladen, Hoke, Robeson, and Scotland Counties, were collected in the fall of 1993. Water levels in the third study unit were collected during the spring of 1994; this unit included the southern parts of Bladen and Robeson Counties. Also, water-level data were collected in a few wells outside of these units to aid in the regional assessment of water-level conditions.

#### PHYSIOGRAPHIC SETTING

The study area is in the Coastal Plain Province of southeastern North Carolina. Stuckey (1965) recognized two physiographic subdivisions in the northern Coastal Plain in this area—the Inner Coastal Plain and the Sand Hills (fig. 1). The Inner Coastal Plain is in the eastern part of the study area and includes most of Bladen and Robeson Counties and parts of Hoke and Scotland Counties. This area consists of broad, flat, continuous swampy uplands between major streams. Local topographic relief is generally 5 feet (ft) or less. However, land near the major rivers, such as the Cape Fear and its tributaries, is quite dissected. In some areas, these streams may be incised 20 ft or more into the flat uplands.

The Sand Hills area is in the western part of the study area and includes the northern two-thirds of Hoke County and the northwestern half of Scotland County (Schipf, 1961). This area is characterized by rolling hills, deep sand, and sandy soils (Stuckey, 1965), upon which a dendritic drainage pattern has developed. The larger stream of the area flows eastward or southeastward across the Coastal Plain, and stream valleys are characterized by steep sides and well-developed flood plains.

#### HYDROGEOLOGIC SETTING

The Coastal Plain Province of North Carolina consists of a southeastward dipping and thickening wedge of predominantly unconsolidated sedimentary deposits of sand and clay underlain by basement rocks. The western boundary of these deposits is the Fall Line (fig. 1). The sedimentary deposits are divided into geologic formations based on their age and lithology. The deposits are further classified according to hydraulic and geologic characteristics and grouped into hydrogeologic units called aquifers and confining units. The basement rocks consist of metamorphic and igneous crystalline rocks of pre-Cretaceous age. The correlation of geologic and hydrogeologic units that underlie the study area is shown in figure 3.

Aquifers are composed of a formation, parts of formations, or groups of formations that (1) contain ample saturated permeable material (sand or gravel in this study area), (2) allow the lateral movement of water within them, and (3) yield significant quantities of water to wells and springs (see table 1) that restrict the movement of water between adjacent aquifers.

Wimmer and Coble (1989) identified six aquifers in the study area—the surficial, Yorktown, Peedee, Black Creek, upper Cape Fear, and lower Cape Fear aquifers—separated by five confining units (figs. 3 and 4). Each confining unit is informally named for the aquifer it overlies. The Yorktown aquifer is not present in most of the study area except in Robeson County near the South Carolina State line.

The surficial aquifer is an unconfined aquifer that consists generally of thin hydrogeologic units near land surface. The water table, which is the upper surface of the water in the surficial aquifer, is at atmospheric pressure. Unconfined conditions may occur in the Yorktown, Peedee, and Black Creek aquifers where they are at shallow depths and where their respective confining units are missing. An aquifer that occurs beneath confining units is called a confined aquifer. The ground water in a confined aquifer is at a pressure greater than atmospheric.

Where present, the Black Creek confining unit overlies the Black Creek aquifer and is composed of clay, silt-clay, and sandy-clay beds (Wimmer and Coble, 1989). It ranges in thickness from zero to 50 ft.

The Black Creek aquifer is the major source of water for public, industrial, and agricultural uses in the study area and is emphasized in this report. The Black Creek aquifer ranges in thickness from about 100 to 250 ft and consists mainly of sediments of both the Black Creek and the underlying Middleford Formations of Cretaceous age (Wimmer and Coble, 1989). The Black Creek Formation is composed of thinly laminated gray to black clay interlayered with gray to tan sand. The lagoonal or marine origin of the Black Creek Formation is reflected by the shell and organic material, such as lignitized wood, common in these sediments. The Middleford Formation, which outcrops to the west of the Black Creek Formation in the Sand Hills area, is composed of fluvial sediments of nonmarine origin; these include fine- to medium-grained sand beds, interlayered with silt-clay beds and thinly laminated beds of sand and clay.

#### DATA COLLECTION AND ANALYSIS

Wells used during the study included North Carolina Department of Environment, Health, and Natural Resources (DEHNR) observation wells and industrial, municipal, county, and privately owned water-supply wells. To ensure that water levels from water-supply wells closely represented stable conditions in the wells, pumps were turned off at least 20 minutes prior to measuring the water levels. If two consecutive measurements taken about 5 minutes apart differed by more than 0.1 ft, then the water levels were considered representative of nonpumping (quasi-stable) conditions.

Water levels were measured in 56 wells in the fall of 1992 (table 2), 135 wells in the fall of 1993 (table 2), and 62 wells in the spring of 1994 (table 3). Water levels in many of these wells also were measured in the fall of 1988 (Strickland and others, 1992). Of these wells, six are located outside of the LRCOG area—three in Scotland County, North Carolina, and three in Dillon County, South Carolina (fig. 5). Two wells were equipped with water-level recorders set to automatically record water levels at 1-hour intervals, and four wells were measured on a periodic basis to monitor water-level fluctuations in the Black Creek aquifer during the 3-year study (fig. 5).

The water level measured in a well represents the hydraulic head in the screened part of the aquifer. This water level is a function of the elevation of the screened part of the well and the pressure within the aquifer. The water level measured in a well that taps an unconfined aquifer indicates the altitude of the water table in the surrounding aquifer. Below the water table, the pressure is greater than atmospheric and increases with increasing depth. In confined aquifers, which are generally deeper aquifers, the altitude of water in a tightly cased, nonpumped well will stand above the top of the aquifer; this hydraulic head defines a point on the potentiometric surface of an aquifer. A map of the potentiometric surface depicts the areal distribution of hydraulic heads and can be used to infer the direction of ground-water movement in the aquifer. Lines of equal hydraulic head in the aquifer are represented by potentiometric-surface contour lines. Ground-water flow through an aquifer from areas of high hydraulic head to areas of low hydraulic head. The direction of flow is assumed to be perpendicular to the potentiometric-contour lines.

The water-level data were used to construct three maps of the potentiometric surface of the Black Creek aquifer—one for each study unit (fig. 5). Because water-level data were collected during three different time periods, the maps of the potentiometric surface were constructed such that the potentiometric contours do not cross over into adjacent study units. However, there was some overlap of the study units (fig. 5). Water levels in several wells were measured more than once during different data-collection periods (table 4). Where multiple measurements were available, the latest water-level measurement was used to construct the potentiometric surface. Thus, the potentiometric surface configuration of the Black Creek aquifer during the fall of 1992 is slightly different in some areas from that reported by Strickland (1994) because of the additional water-level data.

Well-construction data, including depths of casing and screened intervals, also were collected. By using the Coastal Plain hydrologic framework developed by Wimmer and Coble (1989), each well was assigned a primary aquifer based on the position of the screened interval. Most wells are screened in only one aquifer, but several wells are screened in multiple aquifers (tables 1 and 3). Water levels measured in these wells represent a combination of hydraulic heads in the screened interval. Hydraulic heads in a single aquifer at some distance from these wells are likely different from the composite hydraulic heads.

The locations of many wells, mainly in the second and third study units, were determined by using the Global Positioning System (GPS). This technology, which uses information from satellites to compute geographic positions, can determine latitude and longitude coordinates. Latitude and longitude coordinates obtained by using GPS are listed in tables 1-3 and may be different from those reported by Strickland (1994). Land-surface altitudes at well sites, which had not been surveyed to the nearest one-tenth of a foot, were estimated to the nearest foot from USGS 7.5-minute topographic quadrangle maps having 5- or 10-ft contour intervals.

Hydraulic heads were computed by subtracting the water levels in the wells, in feet below land surface, from land-surface altitudes at the wells, in feet above sea level. These values are reported to the nearest foot (tables 1-3).

Ground-water flow in the Black Creek aquifer is controlled by recharge, discharge to perennial streams, and pumped wells. In areas away from pumping centers, the natural pattern of ground-water flow is not disrupted, and hydraulic heads in nonpumped wells near streams are slightly to substantially higher than the stream stage. Therefore, to construct potentiometric contours, the hydraulic head in the unconfined part of the aquifer along a stream was assumed to be slightly greater than the altitude of the stream.

Withdrawals from an aquifer can modify the natural pattern of ground-water flow. Pumping water from an aquifer can cause hydraulic heads to decline in the vicinity of the wells being pumped. If the hydraulic heads decline, the potentiometric surface of the aquifer is depressed so that it resembles the shape of an inverted cone which has its lowest point at the center of pumping. Such a configuration is referred to as a cone of depression.

#### WATER-LEVEL CONDITIONS IN THE BLACK CREEK AQUIFER

The potentiometric surface of the Black Creek aquifer throughout the mapped area in Hoke, Robeson, and Scotland Counties and parts of Bladen County (fig. 5) shows that ground water flows from areas of high hydraulic head to areas of low hydraulic head. Where potentiometric contours cross streams, they bend upstream, indicating that ground water is discharging from the aquifer into the streams. In the highly dissected Sand Hills area, the aquifer is unconfined in stream valleys where streams have cut through the Black Creek confining unit (Wimmer and Coble, 1989). Evidence for the unconfined condition of the Black Creek aquifer is shown in figure 5 where the orientation of the potentiometric contours is closely spaced and strongly influenced by the topography of the streams. Also, farther east in Bladen County, the channel of the Cape Fear River between the Cumberland-Bladen County line and Elizabethtown is cut through the Black Creek confining unit to allow direct hydraulic connection between the river and the Black Creek aquifer (Wimmer and Coble, 1989). Evidence of this hydraulic connection is shown by the closely spaced potentiometric contours along the Cape Fear River.

The natural flow is disrupted around pumping centers where ground water is discharged from pumped wells. Withdrawals from wells at pumping centers in the Elizabethtown and Lumberton areas have resulted in substantial cones of depression in the potentiometric surface of the Black Creek aquifer. In the fall of 1992, the lengths of the major axes of the cones of depression were about 4 and 6 miles (mi) long beneath the Elizabethtown and Lumberton areas, respectively (Strickland, 1994).

At a pumping center in western Robeson County, withdrawals from the Black Creek aquifer 3 mi north of Maxton (area A, fig. 5) averaged 5 to 6 million gallons per day (Mgal/d) during the fall of 1993 (Mr. Myron Neville, oral communication, 1995); however, the cone of depression that developed around this well field was not extensive in the fall of 1993. The effects of pumping on the potentiometric surface at other local pumping centers were not mappable at 10- or 20-ft contour intervals. These included pumping centers at Lumberton (area C, fig. 5) and east of Laurinburg (area B, fig. 5) in Scotland County, and Raeford (area D, fig. 5) in Hoke County. Withdrawals at these pumping centers averaged 2.3 Mgal/d (Mr. Robert Ellis, oral communication, 1995), 1.1 Mgal/d (Mr. Gary Arnett, oral communication, 1995), and 1.5 Mgal/d (Mr. Jimmy Allen, oral communication, 1995), respectively, during 1993.

In the mapped part of southwestern Bladen County (fig. 5), the Black Creek aquifer is overlain by the Peedee aquifer and the Black Creek confining unit. Most water in this part of the Black Creek aquifer does not discharge to local streams but flows southeastward toward the coast or is discharged from pumped wells. Withdrawals from wells at pumping centers, such as those in Clarkson (fig. 5) where the observed water level declined 1 ft between the fall of 1988 and the spring of 1994 (well BL-22, table 2). This change could have been a seasonal response with little or no influence from pumping.

During the fall of 1993 and the spring of 1994, observed water levels in the Black Creek aquifer at or near pumping centers in southern Bladen, Hoke, western Robeson, and Scotland Counties were similar or slightly higher than those measured in the fall of 1988 (fig. 5). An example of this is observation well HO-47, which is near water-supply wells in and around the northwestern part of Raeford in Hoke County. The observed water level at this site was 0.3 ft higher in October 1993 than in December 1988 (table 5). The hydrograph for well HO-47 (fig. 6) indicates that the water level in the Black Creek aquifer responded to withdrawals for water supply and seasonal climatic effects from October 1993 through November 1994. The steep water-level declines and rises of about 1 to 2 ft that occurred over time periods of about 1 week were caused by pumping at nearby wells. The more long-term decline and rise of water levels of about 2 to 3 ft represents seasonal fluctuations. Ground-water-level declines are caused by discharge, both natural and from pumping, that exceeds recharge to the aquifer, and water-level rises are caused by recharge that exceeds discharge. The observed water-level trend at observation well HO-47 from October 1993 through November 1994 indicates that recharge to the Black Creek aquifer exceeded discharge.

In most areas away from pumping centers, observed water-level changes have been minor since 1988 (fig. 5, table 5). Hydrographs for observation wells CO-125 and RB-168 (fig. 6) indicate that water-level fluctuations are caused largely by seasonal variation in ground-water recharge.

The hydrograph for observation well RB-185 (fig. 6), which is located approximately 4 mi east of Lumberton, shows considerable water-level fluctuations in the Black Creek aquifer from September 1992 through December 1994. These fluctuations appear to be the result of nearby pumping and seasonal variation. During this period of record, water-level declines of about 11 ft were recorded between May 1993 and December 1994. The greatest decline, about 9 ft, occurred from May 1993 through November 1993.

In northwestern Bladen County near Tar Heel, the observed water level in well BL-153 declined about 5 ft between September 1992 and June 1993, and then remained fairly steady through October 1994. Major withdrawals, which began in mid-October 1992, from the aquifer below the Black Creek, the upper Cape Fear aquifer, have increased to approximately 2.3 Mgal/d in 1994 in the Tar Heel area (Strickland, 1995). The observed water level in the upper Cape Fear aquifer near the center of pumping declined 50 ft between September 1992 and October 1994 in observation well BL-142, which is located about 1,300 ft northwest of well BL-153 (figs. 5 and 7). The data suggest that this pumping resulted in a large vertical hydraulic gradient between the aquifers near the pumping center and induced leakage of ground water from the Black Creek aquifer through the confining unit into the upper Cape Fear aquifer. This leakage caused the lowering of the hydraulic head in the Black Creek aquifer and, therefore, in the water level in well BL-153. This interpretation is supported by the fact that the large decline in water level in the BL-153. This water-level decline is well BL-153 (fig. 7). Also, between June 1993 and October 1994, the relatively steady water-level decline in the upper Cape Fear aquifer coincided with a fairly steady water level in well BL-153.

At Rowland in southwestern Robeson County, the observed water level in the Black Creek aquifer declined 1.3 ft in well RB-84 between the fall of 1988 and the spring of 1994 (table 5). However, the water level in observation well RB-148, which is located about 1 mi south-southeast of well RB-84, declined 5.4 ft during this same period (table 5). Observed water-level declines in well RB-148, which declined 11.1 ft between July 1981 and October 1994 (fig. 6), are likely a result of pumping by several wells in South Carolina. These withdrawals have caused a regional cone of depression to develop that has extended into North Carolina near Rowland (fig. 5).

#### REFERENCES CITED

- North Carolina Department of Natural Resources and Community Development, 1985, Geologic map of North Carolina, 1985: North Carolina Department of Natural Resources and Community Development, 1 sheet, scale 1:500,000.
- Schipf, R.G., 1961, Geology and ground-water resources of the Fayetteville area, North Carolina: Department of Water Resources, Division of Ground Water, Groundwater Bulletin No. 3, 99 p.
- Strickland, A.G., 1994, Water-level conditions in the Black Creek and upper Cape Fear aquifers, 1992, in parts of Bladen and Robeson Counties, North Carolina: U.S. Geological Survey Water-Resources Investigations Report 94-4016, 1 sheet.
- Strickland, A.G., 1995, Water-level conditions in the upper Cape Fear aquifer, 1992-94, in parts of Bladen and Robeson Counties, North Carolina: U.S. Geological Survey Water-Resources Investigations Report 95-4129, 1 sheet.
- Strickland, A.G., Coble, B.W., Edwards, L.A., and Pope, B.F., 1992, Ground-water-level data for North Carolina, 1988-90: U.S. Geological Survey Report 92-57, 167 p.
- Stuckey, J.L., 1965, North Carolina: its geology and mineral resources: Raleigh, North Carolina Department of Conservation and Development, 550 p.
- Wimmer, M.D., Jr., and Coble, B.W., 1989, A hydrogeologic framework of the North Carolina Coastal Plain aquifer system: U.S. Geological Survey Open-File Report 89-060, 155 p.

**Sea Level.** In this report, "sea level" refers to the National Geodetic Vertical Datum of 1929—a geodetic datum derived from a general adjustment of the first-order level nets of the United States and Canada, formerly called Sea Level Datum of 1929.

**Table 1.** Selected records for wells in the Black Creek aquifer in Bladen and Robeson Counties, September-October 1992  
(Well number refers to well location in figure 5; BL, Bladen County; RB, Robeson County)

Well number	Latitude	Longitude	Altitude of land surface (feet above sea level)	Date of water-level measurement	Water level (feet below or above (+) land surface)	Hydraulic head (feet above or below (+) sea level)
BL-56	34°40'45"	78°54'44"	72	9-92	20.9	51
BL-98	34°43'33"	78°47'29"	125	9-92	25.9	99
*BL-121	34°37'26"	78°36'02"	120	9-92	105.8	140
*BL-127	34°38'35"	78°38'36"	130	9-92	180.4	-50
*BL-129	34°38'53"	78°38'43"	135	9-92	185.6	-51
BL-132	34°45'24"	78°48'24"	130	9-92	61.2	69
BL-143	34°45'03"	78°44'09"	84	9-92	21	63
BL-146	34°38'23"	78°44'22"	139.8	9-92	24.7	115
*BL-149	34°36'52"	78°36'32"	105	9-92	67.2	38
BL-151	34°45'39"	78°48'12"	142	9-92	81.3	61
BL-153	34°44'30"	78°48'18"	128	9-92	50.1	78
BL-154	34°48'57"	78°50'33"	128	10-92	22.6	105
BL-160	34°41'04"	78°47'47"	131	10-92	21.5	110
BL-162	34°49'11"	78°48'06"	75	10-92	35.1	40
BL-167	34°37'53"	78°43'13"	128	10-92	16.1	112
BL-168	34°39'35"	78°45'39"	135	9-92	14.7	120
BL-169	34°42'45"	78°48'55"	135	9-92	15.7	120
BL-171	34°43'06"	78°45'06"	150	3-93	106.1	44
RB-90	34°48'17"	78°58'21"	165	10-92	17.9	147
RB-92	34°48'30"	78°59'07"	175	10-92	22.7	152
RB-97	34°35'31"	79°00'11"	117	10-92	63.5	54
RB-98	34°35'12"	79°00'09"	112	10-92	65.5	47
RB-100	34°35'28"	79°00'05"	117	10-92	65.2	52
RB-104	34°37'24"	79°04'32"	126.6	10-92	6.6	120
BL-113	34°39'05"	78°55'13"	150	10-92	42.3	108
BL-115	34°33'56"	78°59'39"	114	10-92	54.4	60
BL-117	34°47'55"	78°56'17"	158	10-92	18.4	140
BL-130	34°46'08"	79°05'50"	129.6	10-92	14.9	140
BL-157	34°42'36"	79°00'07"	142	9-92	29.9	112
BL-185	34°38'40"	78°55'00"	110	10-92	27.6	82
BL-195	34°37'52"	79°01'33"	107	10-92	28.6	78
RB-198	34°37'50"	79°01'33"	117	10-92	17.6	99
RB-200	34°37'53"	79°01'33"	105	10-92	22.2	83
RB-202	34°38'00"	79°01'52"	105	10-92	22.2	83
RB-207	34°39'28"	79°04'14"	122	10-92	-1.1	123
RB-208	34°39'50"	79°05'52"	148	10-92	11.2	137
RB-209	34°39'32"	79°06'31"	152	10-92	15.4	137
BL-210	34°39'22"	79°05'50"	150	10-92	15	135
BL-211	34°42'08"	79°01'99"	150	10-92	9.8	140
BL-213	34°41'19"	79°03'55"	147	10-92	15.4	132
BL-215	34°39'21"	79°03'44"	124	10-92	4.7	119
BL-216	34°40'44"	79°07'26"	158	10-92	7.8	150
BL-217	34°46'10"	78°55'05"	160	10-92	17.8	142
BL-221	34°36'23"	78°58'32"	151	10-92	34.6	116
BL-222	34°44'45"	78°57'29"	150	10-92	8.6	141
BL-226	34°35'17"	78°58'27"	132	10-92	64.5	68
BL-233	34°38'00"	79°02'14"	116	10-92	21.8	94
BL-234	34°37'42"	79°01'31"	112	10-92	33.8	78
BL-236	34°37'25"	79°01'03"	110	10-92	38.9	71
BL-239	34°41'02"	79°03'31"	152	10-92	25	127
BL-240	34°39'27"	78°55'12"	152	10-92	48.2	104
BL-241	34°48'03"	78°54'08"	140	10-92	12.5	128
BL-242	34°48'09"	78°53'44"	133	10-92	19.5	114
BL-243	34°48'19"	78°53'50"	137	10-92	15.8	121
BL-244	34°48'14"	78°53'54"	134	10-92	11.6	122
BL-245	34°36'47"	79°00'02"	114	10-92	51.2	63

\*Well screened in the Black Creek, the upper Cape Fear, and the lower Cape Fear aquifers.

\*Well screened in the Black Creek and the upper Cape Fear aquifers.

**Table 2.** Selected records for wells in the Black Creek aquifer in Hoke, Robeson, and Scotland Counties, October-December 1993  
(Well number refers to well location in figure 5; HO, Hoke County; RB, Robeson County; SC, Scotland County)

Well number	Latitude	Longitude	Altitude of land surface (feet above sea level)	Date of water-level measurement	Water level (feet below or above (+) land surface)	Hydraulic head (feet above or below (+) sea level)
HO-22	34°58'57"	79°14'01"	258	11-93	34.5	224
HO-32	35°03'17"	79°21'35"	350	12-93	3.2	347
HO-35	34°58'20"	79°14'12"	251	11-93	27.4	224
HO-36	34°59'00"	79°14'16"	276	11-93	52.3	224
HO-37	34°58'07"	79°13'42"	248	11-93	31.1	217
HO-38	34°58'15"	79°13'48"	248	11-93	29.6	218
HO-40	34°59'15"	79°14'43"	290	11-93	66.3	224
HO-47	34°59'34"	79°14'42"	274.4	10-93	42.5	232
HO-48	34°56'52"	79°07'40"	257	11-93	30.8	204
HO-51	34°59'03"	79°08'08"	295.2	11-93	41.5	236
HO-52	34°58'41"	79°15'35"	273.5	11-93	41.5	236
HO-53	34°58'40"	79°16'07"	286.3	11-93	49.6	237
HO-54	34°52'39"	79°17'07"	241	12-93	11	230
HO-55	34°57'27"	79°02'24"	175	09-93	33.8	206
HO-60	34°58'58"	79°16'49"	296.2	11-93	49.5	221
HO-61	34°59'08"	79°16'49"	296.2	11-93	27.2	269
HO-62	34°58'46"	79°16'55"	284	11-93	30.6	253
HO-64	34°58'18"	79°15'26"	268	11-93	24.8	244
HO-65	34°59'11"	79°14'21"	281	11-93	51.9	229
HO-66	34°59'39"	79°14'15"	190	10-93	20	170
HO-70	34°59'40"	79°19'28"	354	12-93	63	291
HO-72	34°52'33"	79°13'19"	225	12-93	16.6	208
HO-75	34°56'43"	79°14'55"	233	12-93	15.3	218
HO-77	34°56'09"	79°20'43"	265	12-93	20.3	225
HO-78	34°58'33"	79°06'35"	241	12-93	31.6	214
HO-79	34°58'36"	79°11'04"	223	12-93	33.9	185
HO-79	34°58'36"	79°11'04"	218	12-93	13.8	204
HO-80	34°59'33"	79°12'14"	234	10-93	43.3	191
HO-81	34°59'28"	79°11'56"	237	10-93	44.8	190
HO-82	34°56'37"	79°15'17"	232	10-93	5	227
HO-83	34°57'21"	79°12'55"	190	11-93	45.1	143
HO-84	34°57'57"	79°20'24"	254	12-93	22	235
HO-85	34°57'56"	79°21'06"	240	12-93	8.9	233
HO-86	34°59'14"	79°12'30"	316	12-93	20.3	288
HO-87	34°58'37"	79°06'35"	241	10-93	38.8	184
HO-88	34°59'08"	79°16'11"	210	10-93	24	189
HO-89	34°54'15"	79°12'40"	225	10-93	14	211
HO-90	35°01'11"	79°03'50"	175	11-93	4.9	177
HO-92	35°01'16"	79°09'19"	250	10-93	20.3	239
HO-93	35°01'09"	79°08'11"	242	11-93	28.2	21
HO-95	35°01'12"	79°06'54"	250	10-93	36	211
HO-96	35°00'04"	79°03'33"	242	10-93	47.9	199
HO-97	35°02'04"	79°06'48"	271	10-93	42.7	222
HO-100	35°02'49"	79°07'10"	274	10-93	39.1	233
HO-102	34°52'52"	79°04'02"	200	10-93	16	218
HO-103	35°03'55"	79°06'35"	241	10-93	33.3	233
HO-104	34°59'18"	79°12'02"	242	10-93	35.6	203
HO-106	34°52'57"	79°14'00"	231	12-93	14.6	211
HO-107	35°02'12"	79°18'03"	347	12-93	54.4	295
HO-108	35°01'43"	79°22'16"	291	12-93	14.2	277
HO-109	34°59'41"	79°05'49"	224	10-93	27	159
HO-110	35°02'19"	79°06'56"	273	10-93	43.8	222
HO-111	34°52'53"	79°12'30"	233	10-93	24.3	202
HO-113	35°00'55"	79°11'16"	283	10-93	43.1	242
HO-114	35°02'10"	79°06'45"	262	10-93	25.5	233
HO-115	34°57'18"	79°03'55"	130	10-93	5.8	127
HO-116	34°58'58"	79°11'01"	272	10-93	55.2	217
HO-121	34°59'50"	79°11'01"	259	11-93	30.3	222
HO-122	34°59'02"	79°10'48"	236	12-93	33	202
HO-123	34°59'11"	79°11'31"	245	12-93	40.6	206
HO-125	35°02'40"	79°08'20"	251	11-93	35	212
HO-126	34°58'33"	79°07'06"	173	12-93	18	159
HO-128	34°58'37"	79°07'11"	203	12-93	45.5	143
HO-129	34°57'00"	79°02'33"	136	11-93	5.9	121
RB-87	34°40'52"	79°15'42"	172	12-93	12.8	181
RB-106	34°39'08"	79°12'34"	166	12-93	18.4	167
RB-108	34°40'52"	79°12'34"	190	12-93	16.5	165
RB-111	34°52'23"	78°58'40"	180	12-93	28.2	101
RB-121	34°43'39"	79°10'56"	167	12-93	9.3	161
RB-131	34°52'53"	79°06'47"	204	12-93	11.2	191
RB-134	34°48'45"	79°12'38"	205	10-93	14	161
RB-165	34°52'18"	79°05'52"	201	12-93	12.6	161
RB-168	34°50'55"	79°05'16"	187.3	10-93	6.4	161
RB-174	34°54'29"	79°02'24"	184	12-93	5.9	161
RB-205	34°46'52"	79°18'23"	212	11-93	19.5	119
RB-230	34°46'28"	79°18'01"	203	11-93	18.5	119
RB-246	34°46'08"	79°19'45"	202	11-93	27.2	119
RB-247	34°46'42"	79°19'37"	192	11-93	26.3	119
RB-248	34°46'57"	79°19'37"	193	11-93	26.3	119
RB-249	34°47'07"	79°19'47"	194	11-93	26.2	119
RB-251	34°52'39"	79°05'22"	200	12-93	14.6	104
RB-252	34°52'00"	79°06'14"	186	12-93	3	104
RB-253	34°53'13"	79°03'54"	178	12-93	10.4	104
RB-255	34°48'22"	79°11'36"	204	10-93	20.8	104
RB-256	34°48'24"	79°11'48"	203	10-93	19.6	104
RB-260	34°40'00"	79°12'12"	165	12-93	15.8	118
RB-262	34°41'34"	79°15'22"	160	12-93	12.7	118
RB-263	34°46'21"	79°19'24"	195	12-93	10	118
RB-265	34°56'50"	79°02'33"	144	11-93	8.2	104
RB-303	34°42'21"	79°21'08"	196	11-93	19.7	217
RB-340	34°45'17"	79°20'35"	210	10-93	21.7	203
SC-61	34°53'15"	79°21'40"	227	12-93	10.8	162
SC-63	34°53'32"	79°26'45"	275	12-93	32.3	162
SC-64	34°54'30"	79°25'01"	276	12-93	10.1	162
SC-65	34°44'28"	79°30'41"	190	12-93	9.2	162
SC-68	34°45'21"	79°27'35"	215	12-93	16.5	162
SC-69	34°44'39"	79°28'03"	212	12-93	16.3	162
SC-70	34°45'37"	79°28'03"	210	11-93	2.7	162
SC-73	34°45'39"	79°28'35"	232	11-93	26.9	169
SC-74	34°46'41"	79°25'10"	210	12-93	16.4	169
SC-77	34°47'02"	79°26'40"	187	11-93	11.1	169
SC-85	34°44'49"	79°28'16"	206	12-93	20.7	169
SC-86	34°44'33"	79°28'32"	205	12-93	23.2	169
SC-87	34°44'18"	79°28'47"	211	12-93	21.3	169
SC-88	34°44'08"	79°29'12"	207	12-93	13.8	169
SC-89	34°44'51"	79°28'39"	205	12-93	17.5	169
SC-92	34°46'13"	79°29'47"	205	12-93	21	169
SC-94	34°43'40"	79°30'27"	215	12-93	23.5	169
SC-96	34°43'20"	79°31'28"	250	12-93	36.8	169
SC-98	34°41'23"	79°31'41"	200	12-93	23.9	169
SC-97	34°49'13"	79°31'49"	200	12-93	5.5	169
SC-98	34°41'50"	79°29'08"	195	12-93	14.7	169
SC-99	34°49'13"	79°25'08"	197	12-93	2.9	169
SC-100	34°52'06"	79°24'27"	230	12-93	49	175
SC-101	34°52'30"	79°25'17"	250	12-93	17.5	175
SC-102	34°51'33"	79°27'47"	365	12-93	53	175
SC-104	34°57'23"	79°20'11"	305	12-93	44	175
SC-105	34°58'07"	79°20'09"	250	12-93	20.4	175
SC-106	34°53'13"	79°22'09"	235	12-93	14.8	175
SC-107	34°53'12"	79°22'11"	235	12-93	15	175
SC-108	34°46'43"	79°29'12"	230	12-93	9.8	175
SC-110	34°48'19"	79°23'10"	220	12-93	24.9	175
SC-111	34°49'31"	79°24'42"	230	12-93	27.9	175
SC-112	34°52'29"	79°20'47"	220	12-93	11.6	175
SC-113	34°52'08"	79°21'06"	221	12-93	8.6	175
SC-115	34°42'36"	79°33'09"	210	12-93	18.5	175
SC-116	34°48'35"	79°31'31"	241	12-93	17.9	175
SC-117	34°46'48"	79°32'43"	227	12-93	28.8	175
SC-118	34°46'20"	79°32'43"	255	12-93	14.3	175
SC-119	34°47'17"	79°32'17"	211	12-93	14.3	175
SC-121	34°46'00"	79°22'12"	212	12-93	32	175
SC-123	34°46'03"	79°22'25"	198	12-93	16.4	175
SC-126	34°46'14"	79°22'29"	186	12-93	16	175