

## INTRODUCTION

The Highland Lakes on the Colorado River are in an area periodically threatened by large storms and floods. Many storms exceeding 10 inches (in) in depth have been documented in the area, including some that peaked approaching 40 in. These storms typically produce large peak discharges that often threaten lives and property. The storms sometimes occur with little warning. Steep stream slopes and thin soil characteristics of the area often cause large peak discharges and rapid movement of floods through watersheds. A procedure to predict the discharge associated with large floods is needed for the area at that appropriate peak discharges can be used in the design of flood plains, bridges, and other structures.

The U.S. Geological Survey (USGS), in cooperation with the Lower Colorado River Authority (LCRA), studied flood peaks for streams in the vicinity of the Highland Lakes of central Texas. The Highland Lakes are a series of reservoirs constructed on the Colorado River. The chain of lakes (and year each was completed) comprises Lake Buchanan (1937), Lake Llynnon (1938), Lake Lyndon B. Johnson (1950), Lake Marble Falls (1951), Lake Travis (1942), and Lake Austin (1980). The study area (fig. 1), which includes all or parts of 21 counties in the vicinity of the Highland Lakes, was selected because most streams in the area have flood characteristics similar to streams entering the Highland Lakes. The entire study area is in a region subject to large storms.

The purpose of this report is to present (1) peak-flow frequency data for stations and equations to estimate peak-flow frequency for large streams with natural drainage basins in the vicinity of the Highland Lakes, and (2) a technique to estimate the extreme flood peak discharges for the large streams in the vicinity of the Highland Lakes. Peak-flow frequency in this report refers to the peak discharges for recurrence intervals of 2, 5, 10, 25, 50, and 100 years. A large stream is defined as having a contributing drainage area of at least 0.5 square mile (mi<sup>2</sup>); and a natural drainage basin has less than 10 percent impervious cover and less than 10 percent of its drainage area controlled by reservoirs.

The mean annual precipitation in the study area for 1951–80 ranges from about 20 in. in western Kimble County to about 34 in. at the eastern edge of Williamson County (Riggin and others, 1987, p. 23). Many large storms and catastrophic floods have occurred along or in the adjacent area west of the Balcones escarpment (fig. 1). (Daly and others, 1987, p. 23; Breeding and Daly, 1964; Breeding and Montgomery, 1954; Schroeder and others, 1979; Caran and Baker, 1986; Slade, 1986; and Hegi and others, 1986). About a dozen storms with precipitation depths exceeding 15 in. in a few days or less have been documented in this area during the past 60 years. Some of these storms have produced world-record precipitation depths for durations less than 48 hours. The documentation for these storms and for other large storms indicates that they are not uniformly distributed temporally or spatially; therefore, the recurrence intervals for such storms cannot be varied (Slade, 1986, p. 17). These large storms can cause flood peaks that would exceed those that can be predicted accurately by analyses of available precipitation or flood data.

The peak-flow frequency was estimated for each of 55 qualified stations in the study area (table 1) following guidelines established by the Interagency Advisory Committee on Water Data (1982). Qualified streamflow-gaging stations for the study area are those with at least 8 years of data from natural drainage basins (sites 1–55, fig. 1). Equations to estimate peak-flow frequency for large streams with natural drainage basins in the vicinity of the Highland Lakes were developed. These equations were developed from selected stations on the basis of the relationship between peak-flow frequency and basin characteristics for each station. The entire period of systematic record (through 1993) was used in the frequency analyses for each qualified station except for stations at which streamflow was regulated during part of the record. These stations are Leon River near Belton (site 1), Lampasas River near Youngsburg (site 5); North Fork San Gabriel River near Georgetown (site 6); San Gabriel River at Lampert (site 12); Brady Creek at Brady (site 16); San Saba River at San Saba (site 18); Rebecca Creek near Spring Branch (site 41); and Cibola Creek near Boerne (site 54). One or more reservoirs were completed in the basin of each of these stations during the period of systematic record. These reservoirs caused the annual peak discharges to become regulated. The annual peak discharges for 1994 and 1995 at Sandy Creek near Kingsland (site 28) were used to include data associated with extreme flooding that occurred in 1995.

The extreme flood potential in the study area was investigated using an "envelope" or "extreme flood potential" curve. This curve is based on the relation between the contributing drainage area and (1) the maximum peak discharge of record for each qualified station (table 1); (2) substantial peak discharges documented for 84 sites without stations (sites 56–139, fig. 1, table 2); and (3) 100-year peak discharges from peak-flow frequency for stations (table 1). Peak discharges estimated from this curve represent the extreme flood potential for the study area.

## PEAK-FLOW FREQUENCY

### Peak-Flow Frequency for Gaging Stations

Peak discharges are monitored at each of the qualified stations in the study area. These stations have various periods of systematic record, as identified in table 1. The peak discharges used in this investigation include the largest peak discharge for each year of systematic record (annual peak discharge) and all known historical peak discharges through 1993. A historical peak discharge—documented by newspaper articles, personal recollections, or other historical sources—represents the largest peak discharge since a known date preceding the beginning of the systematic record. The historical record is the number of years represented by the historical peak. For example, 31 years of systematic record from 1963–93 exist for the station South Fork Rocky Creek near Briggs (site 4). However, the 1976 peak discharge is the highest since 1904 or before, according

to a local resident; thus that peak is the largest in at least 90 years (1904–93). Historical peak discharges can occur before or within the systematic record.

The annual and historical peak discharges for each station were used, together with the USGS computer program PEAKFQ (Slade and Asquith, 1996), to estimate peak discharges for the 2-, 5-, 10-, 25-, 50-, and 100-year recurrence intervals (table 1). The computer program for peak-flow frequency analysis follows the guidelines established by the Interagency Advisory Committee on Water Data (IACWD) and uses the log-Pearson Type III (LPIII) frequency distribution. An equation of the relation between annual peak discharge and probability of exceedance, with the LPIII frequency curve defined by the data superimposed, for Sandy Creek near Kingsland is shown in figure 2.

The skew in the distribution of annual peaks is characterized by a skew coefficient. A reliable skew coefficient is difficult to estimate for stations having short records. Therefore, the IACWD recommends using a weighted skew coefficient with the LPIII distribution. This weighted skew coefficient is computed by weighting the skew coefficient calculated for a station (station skew coefficient) with a generalized skew coefficient representative of the surrounding area. The weighted skew coefficient is based on the inverse of the respective mean square errors for the station and generalized skew coefficients. Generalized skew coefficients were determined for each station (Linda Judd, U.S. Geological Survey, written commun., 1994; Asquith and Slade, 1995b). A weighted skew coefficient then was used in the computation of the peak-flow frequency for each station except San Gabriel River at Lampert (site 12), North Llano River near Junction (site 19), Llano River near Junction (site 20), Llano River at Llano (site 27), and Rebecca Creek near Spring Branch (site 51). For these stations, only the station skew coefficient was used because it produces a better visual fit of the LPIII frequency curve to the data.

Additionally, the IACWD provides a procedure for estimating low-outlier thresholds; annual peak discharges less than this threshold—low outliers (fig. 2)—are excluded from the fitting of the LPIII frequency curve. The proper estimation of low-outlier thresholds is critical in preventing computations of erroneous peak-flow frequencies. The IACWD procedure for estimating low-outlier thresholds is not always appropriate for stations in Texas. An equation to estimate low-outlier thresholds for stations in Texas has been developed (Asquith and others, 1995). The equation was used to estimate most low-outlier thresholds for this investigation. The equation estimates low-outlier thresholds by using the mean, standard deviation, and skew of the logarithms for the systematic annual peak discharges. At five stations, low-outlier thresholds were identified by visually fitting the LPIII frequency curve to the peak-flow data rather than using the equation—South Fork Rocky Creek near Briggs (site 4); North Llano River near Junction (site 19); Llano River near Junction (site 20); Barton Creek at Loop 360 at Austin (site 38); and Rebecca Creek near Spring Branch (site 51).

**Selected Basin Characteristics for Streamflow-Gaging Stations**  
Selected basin characteristics were aggregated for each station. Only those characteristics considered pertinent for estimation of peak-flow frequency from previous similar investigations in Texas (Schroeder and Massey, 1977; Slade and others, 1995) are included in this investigation. The characteristics are: 2-year 24-hour precipitation, mean annual precipitation, contributing drainage area, stream length, shape factor, and stream slope (table 1). The 2-year 24-hour precipitation and mean annual precipitation are for the period 1951–80 and are expressed in inches. The 24-hour precipitation is the maximum 24-hour precipitation depth exceeding 15 in. in a few days or less have been documented in this area during the past 60 years. Some of these storms have produced world-record precipitation depths for durations less than 48 hours. The documentation for these storms and for other large storms indicates that they are not uniformly distributed temporally or spatially; therefore, the recurrence intervals for such storms cannot be varied (Slade, 1986, p. 17). These large storms can cause flood peaks that would exceed those that can be predicted accurately by analyses of available precipitation or flood data.

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The extreme flood potential in the study area was investigated using an "envelope" or "extreme flood potential" curve. This curve is based on the relation between the contributing drainage area and (1) the maximum peak discharge of record for each qualified station (table 1); (2) substantial peak discharges documented for 84 sites without stations (sites 56–139, fig. 1, table 2); and (3) 100-year peak discharges from peak-flow frequency for stations (table 1). Peak discharges estimated from this curve represent the extreme flood potential for the study area.

**Multiple-Regression Equations to Estimate Peak-Flow Frequency**  
A weighted-least-squares multiple-regression analysis was done to obtain equations to estimate peak discharges for selected recurrence intervals from basin characteristics. The 2-, 5-, 10-, 25-, 50-, and 100-year peak discharges obtained using the IACWD guidelines were the dependent variables, and the basin characteristics were the independent, or explanatory, variables. The dependent and independent variables, except for the 2-year 24-hour precipitation and mean annual precipitation, were transformed to their common base 10 logarithms before the analysis. A weighted-least-squares regression was used because the peak-flow frequency analyses for stations with short record are subject to greater error than those with long record. Empirical equations (G.D. Tisdler, U.S. Geological Survey, written commun., 1994) were used to calculate a record-length weight factor for each station based on the number of years of systematic record, the number of historical peaks, and the number of years of historical record (table 1).

Various sets of stations were tested in the regression analysis to determine the set of stations that provides the best relation between peak discharge and basin characteristics—that is, the set of stations that produced regression equations that provide the most reliable estimates of peak discharge for the selected recurrence intervals. The sets of stations tested were: (1) all 55 stations in the study area; (2) all stations in basins draining directly to the Highland Lakes, predominantly those in the Llano and Pedernales River Basins; (3) all stations in basins that drain directly to the Highland Lakes, plus those stations in the San Saba River-Brady Creek Basins; and (4) all stations within a radius of about 45 miles (mi) from the length.

approximate center of the Highland Lakes area (the town of Marble Falls). The set of stations that provides the best relation between peak discharge and basin characteristics, and thus is used in the analysis, is the set within about 45 mi of Marble Falls. Thirty-four stations in the study area are within about 45 mi of Marble Falls. To minimize the effects of inter-station correlation on the regression analysis, correlation between annual peak discharges for each pair of stations was computed. For those stations in which the coefficient of determination ( $R^2$ ) between data from two stations was greater than or equal to 0.50, the station with the shorter period of record was excluded from the regression analysis. On that basis, three stations were excluded: North Fork San Gabriel River near Georgetown (6), South Fork San Gabriel River near Bertram (7), and Blanco River near Kyle (site 53). Therefore, 31 stations (table 1) were used to develop the regression equations.

Equations were developed using the explanatory variable "stream length" was excluded from the regression analysis because it is highly correlated with the variable "contributing drainage area." Many equations were developed for each recurrence interval using various combinations of the five remaining independent variables (2-year 24-hour precipitation, mean annual precipitation, contributing drainage area, shape factor, and stream slope) in a forward-stepwise regression procedure. Contributing drainage area and stream slope consistently were the most significant explanatory variables. Additional basin characteristics were significant for some recurrence intervals, but they were not retained in the equations. Testing of equations showed that, if the explanatory variables are not consistent among the equations, some combinations of variable values can produce estimated peak discharges for a given recurrence interval that are less than the peak discharges computed for a smaller recurrence interval. Also, keeping only explanatory variables that are significant at all recurrence intervals has little effect on the other coefficient of determination or standard error of estimation. Thus, only contributing drainage area and stream slope appear in the equations shown below.

| Equation for indicated T-year peak discharge (cubic feet per second) | Coefficient of determination ( $R^2$ ) | Standard error of estimation (percent) |
|--|--|--|
| $Q_2 = 6.67(CDA)^{0.089}(SL)^{0.245}$                                | 0.95                                   | 33                                     |
| $Q_5 = 27.0(CDA)^{0.089}(SL)^{0.245}$                                | 0.92                                   | 31                                     |
| $Q_{10} = 15.5(CDA)^{0.089}(SL)^{0.245}$                             | 0.95                                   | 33                                     |
| $Q_{25} = 8.08(CDA)^{0.089}(SL)^{0.245}$                             | 0.92                                   | 32                                     |
| $Q_{50} = 5.13(CDA)^{0.089}(SL)^{0.245}$                             | 0.92                                   | 43                                     |
| $Q_{100} = 3.34(CDA)^{0.089}(SL)^{0.245}$                            | 0.91                                   | 46                                     |

where  
 $Q_T$  = discharge, for specified (T-year) recurrence interval, in cubic feet per second;  
CDA = contributing drainage area, in square miles; and  
SL = stream slope, in feet per mile.

The equations are based on stations in natural drainage basins within about 45 mi of Marble Falls and are applicable for estimating peak discharges for streams in that region. In the process of selecting stations used to develop the regression equations, correlations between peak discharge and basin characteristics were determined for stations as far as 90 mi from Marble Falls. Correlation generally was less for stations 45 to 90 mi from Marble Falls than for stations within 45 mi of Marble Falls. The applicability of the equations to sites farther than about 45 mi from Marble Falls probably decreases with increasing distance from Marble Falls, but determination of a specific distance from Marble Falls beyond which the equations are not applicable is problematic.

The equations might not be applicable for stream sites in basins with contributing drainage area, shape factor, or channel slope not in or outside the ranges of values of the stations used to develop the equations (table 1). In the authors' judgment, the equations can be used for sites with contributing drainage area between about 0.5 and 3,000 mi<sup>2</sup>; for sites with shape factors between about 1.0 and 15; and for sites with stream slopes between about 9 and 100 feet per mile.

For sites at or near the stations included in the analysis, the regression equations might yield estimates of discharge different from those computed using the IACWD guidelines. The dependent variables, and the basin characteristics were the independent, or explanatory, variables. The dependent and independent variables, except for the 2-year 24-hour precipitation and mean annual precipitation, were transformed to their common base 10 logarithms before the analysis. A weighted-least-squares regression was used because the peak-flow frequency analyses for stations with short record are subject to greater error than those with long record. Empirical equations (G.D. Tisdler, U.S. Geological Survey, written commun., 1994) were used to calculate a record-length weight factor for each station based on the number of years of systematic record, the number of historical peaks, and the number of years of historical record (table 1).

Various sets of stations were tested in the regression analysis to determine the set of stations that provides the best relation between peak discharge and basin characteristics—that is, the set of stations that produced regression equations that provide the most reliable estimates of peak discharge for the selected recurrence intervals. The sets of stations tested were: (1) all 55 stations in the study area; (2) all stations in basins draining directly to the Highland Lakes, predominantly those in the Llano and Pedernales River Basins; (3) all stations in basins that drain directly to the Highland Lakes, plus those stations in the San Saba River-Brady Creek Basins; and (4) all stations within a radius of about 45 miles (mi) from the length.

$Q_T = (0.036 \times Q_2) + (MSSE \times Q_2)$  (1)

where  
 $Q_T$  = the weighted 100-year peak discharge, in cubic feet per second;  
 $Q_2$  = the mean square error of the 100-year regression equation;  
 $MSSE$  = the 100-year discharge for the station from an LPIII frequency analysis (table 1), in cubic feet per second;

$MSSE$  = the mean square error of the 100-year discharge for the station (table 1); and  
 $Q_2$  = the 100-year discharge from the regression equation, in cubic feet per second.

The weighted 100-year peak discharge for Sandy Creek near Kingsland (site 28) relative to the LPIII frequency curve and a frequency curve constructed from equation-derived peaks is shown in figure 2. Mean square errors for recurrence intervals other than the 100-year peak discharge are not presented in this report.

## EXTREME FLOOD POTENTIAL

Knowledge of the extreme flood potential for streams in the Highland Lakes area is useful for evaluating extreme floods. To aid in extreme flood evaluation, the USGS has routinely documented substantial peak discharges at stream sites without gaging stations. Many of these substantial peak discharges are associated with catastrophic storms. These sites are identified on figure 1 (sites 56–139), and ancillary information is presented in table 2.

A common technique for investigating extreme flood potential for an area is through the use of an "envelope" or "extreme flood potential" curve (Crippen and Bue, 1977; Thomas and others, 1994; Asquith and Slade, 1995a). This curve is based on the relation between contributing drainage area and peak discharges. Recurrence intervals are not associated with peak discharges from "envelope" curves (Crippen and Bue, 1977). The extreme flood potential curve for the study area (fig. 3) envelopes the maximum peak discharge of record for the stations (table 1), substantial peak discharges for stream sites without stations (table 2), and the 100-year peak discharges for stations in the study area (table 1). As more flood data are collected, peak discharges could exceed those enveloped by the curve, requiring the curve to be redrawn.

The maximum peak discharge of record exceeds the 100-year peak discharge for 17 of the 55 gaging stations. The peak discharges from the regression equation exceed the 100-year peak discharge by at least 30 percent, which indicates that the potential exists for very large floods in the study area.

## SELECTED REFERENCES

Abbot, P.L., and Woodruff, C.M., Jr., eds., 1986, The Balcones escarpment—geology, hydrology, ecology and social development in central Texas: Geological Society of America (GSA) annual meeting, 1986, San Antonio, 200 p.

Asquith, W.H., and Slade, R.M., Jr., 1995a, Documented and potential extreme peak discharges and relation between potential extreme peak discharges and probable maximum peak discharges in Texas: U.S. Geological Survey Water-Resources Investigations Report 95-429, 58 p.

1995b, Flood frequency in Texas—calculation of peak-streamflow frequency at gaging stations: U.S. Geological Survey Fact Sheet 181-95, 2 p.

Asquith, W.H., Slade, R.M., Jr., and Judd, L.J., 1995, Analysis of low-outlier thresholds for log Pearson Type III peak-streamflow frequency analysis in Texas, in Texas Water '95, A Component Conference of the First International Conference on Water Resources Engineering, Water Resources Engineering Division, American Society of Civil Engineers, August 16–17, 1995, San Antonio, Proceedings: San Antonio, American Society of Civil Engineers, p. 379–384.

Baker, V.R., 1975, Flood hazards along the Balcones escarpment in central Texas—alternative approaches to their recognition, mapping, and management: Austin, University of Texas, Bureau of Economic Geology Circular 75-5, 22 p.

Breeding, S.D., and Montgomery, L.H., 1954, Floods of September 1952 in the Colorado and Guadalupe River Basins, central Texas: U.S. Geological Survey Water-Supply Paper 1260-A, 47 p.

Caran, S.C., and Baker, V.R., 1986, Flooding along the Balcones escarpment, central Texas, in Abbott, P.L., and Woodruff, C.M., Jr., eds., The Balcones escarpment—geology, hydrology, ecology and social development in central Texas: Geological Society of America (GSA) annual meeting, 1986, San Antonio, p. 1–14.

Crippen, R.L., and Bue, C.D., 1977, Maximum floodflows in the conterminous United States: U.S. Geological Survey Water-Supply Paper 1887, 51 p.

Daly, C.M., and others, 1987, Major Texas floods of 1935: U.S. Geological Survey Water-Supply Paper 790-4, 68 p.

Hegi, H.R., Slade, R.M., Jr., and Jennings, M.E., 1996, Floods in central Texas, December 1991: U.S. Geological Survey Water-Resources Investigations Report 95-429, 1 sheet.

Interagency Advisory Committee on Water Data, 1982, Guidelines for determining flood flow frequency: Reston, Va., U.S. Geological Survey, Office of Water Data Coordination, Hydrology Subcommittee Bulletin 17B (variously pagel).

Patterson, J.L., 1963, Floods in Texas—magnitude and frequency of peak flows: Texas Water Commission, 31 p.

Riggin, R.F., Bonar, G.W., and Larkin, T.J., 1987, Texas droughts—its recent history (1931–1985): Texas Water Commission Report LP 87-04, 74 p.

Ruggles, F.H., Jr., 1966, Floods on small streams in Texas: U.S. Geological Survey Open-File Report 89, 98 p.

Schroeder, E.E., 1972, Flood stages and discharges for small streams in Texas—compilation of data through September 1970: U.S. Geological Survey Open-File Report, 316 p.

Schroeder, E.E., and Massey, B.C., 1977, Technique for estimating the magnitude and frequency of floods in Texas: U.S. Geological Survey Open-File Report 77-110, 22 p.

Schroeder, E.E., Massey, B.C., and Chih, E.H., 1987, Floods in central Texas, August 1–4, 1978: U.S. Geological Survey Professional Paper 1332, 39 p.

Schroeder, E.E., Massey, B.C., and Waddell, K.M., 1979, Floods in central Texas, August 1978: U.S. Geological Survey Open-File Report 79-682, 121 p.

Slade, R.M., Jr., 1986, Large rainstorms along the Balcones escarpment in central Texas, in Abbott, P.L., and Woodruff, C.M., Jr., eds., The Balcones escarpment—geology, hydrology, ecology and social development in central Texas: Geological Society of America (GSA) annual meeting, 1986, San Antonio, p. 15–20.

Slade, R.M., Jr., and Asquith, W.H., 1996, Peak data for U.S. Geological Survey gaging stations, Texas network, and computer program to estimate peak-streamflow frequency: U.S. Geological Survey Open-File Report 96-148, 57 p.

Slade, R.M., Jr., Asquith, W.H., and Tasker, G.D., 1995, Multiple-regression equations to estimate peak-flow frequency for streams in Hays County, Texas: U.S. Geological Survey Water-Resources Investigations Report 95-409, 1 sheet.

Texas Board of Water Engineers, 1993, Summary of peak flood flow measurements and other measurements of stream discharge in Texas at points other than gaging stations: Texas Board of Water Engineers Bulletin 5807-C, 255 p.

Thomas, B.E., Hjalmarson, H.W., and Walthers, S.D., 1994, Methods for estimating magnitude and frequency of floods in the southwestern United States: U.S. Geological Survey Open-File Report 93-419, 211 p.

Thomas, D.M., 1976, Flood frequency—expected and unexpected probabilities: U.S. Geological Survey Open-File Report 76-715, 7 p.

Williams, G.R., and Crawford, L.C., 1940, Maximum discharges at stream-measurement stations through December 31, 1937, with a supplement including additions and changes through September 30, 1938, by W.S. Eisenhart, Jr.: U.S. Geological Survey Water-Supply Paper 847, 272 p.

Schroeder, E.E., and Massey, B.C., 1977, Technique for estimating the magnitude and frequency of floods in Texas: U.S. Geological Survey Open-File Report 77-110, 22 p.

Schroeder, E.E., Massey, B.C., and Chih, E.H., 1987, Floods in central Texas, August 1–4, 1978: U.S. Geological Survey Professional Paper 1332, 39 p.

Schroeder, E.E., Massey, B.C., and Waddell, K.M., 1979, Floods in central Texas, August 1978: U.S. Geological Survey Open-File Report 79-682, 121 p.

Slade, R.M., Jr., 1986, Large rainstorms along the Balcones escarpment in central Texas, in Abbott, P.L., and Woodruff, C.M., Jr., eds., The Balcones escarpment—geology, hydrology, ecology and social development in central Texas: Geological Society of America (GSA) annual meeting, 1986, San Antonio, p. 15–20.

Slade, R.M., Jr., and Asquith, W.H., 1996, Peak data for U.S. Geological Survey gaging stations, Texas network, and computer program to estimate peak-streamflow frequency: U.S. Geological Survey Open-File Report 96-148, 57 p.

Slade, R.M., Jr., Asquith, W.H., and Tasker, G.D., 1995, Multiple-regression equations to estimate peak-flow frequency for streams in Hays County, Texas: U.S. Geological Survey Water-Resources Investigations Report 95-409, 1 sheet.

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**Table 1.** Selected basin characteristics and peak-flow frequency for streamflow-gaging stations in the vicinity of the Highland Lakes, central Texas.

| fig. 1 | no.      |   |
|--------|----------|---|
| 1      | 08102500 | Leon River near Belton, Texas                       |
| 2      | 08102900 | School Branch near Lampasas, Texas                  |
| 3      | 08103600 | Lampasas River near Kemper, Texas                   |
| 4      | 08103900 | South Fork Rocky Creek near Briggs, Texas           |
| 5      | 08104000 | South Fork Rocky Creek near Yoergroup, Texas        |
| 6      | 08104700 | North Fork San Gabriel River near Georgetown, Texas |
| 7      | 08104850 | South Fork San Gabriel River near Bertram, Texas    |
| 8      | 08104900 | South Fork San Gabriel River at Georgetown, Texas   |
| 9      | 08105000 | San Gabriel River at Georgetown, Texas              |
| 10     | 08105100 | Berry Creek near Georgetown, Texas                  |
| 11     | 08105400 | San Gabriel River near Childress, Texas             |
| 12     | 08105700 | San Gabriel River at Lampert, Texas                 |
| 13     | 08105900 | Avery Branch near Taylor, Texas                     |
| 14     | 08144500 | San Saba River at Abney, Texas                      |
| 15     | 08144600 | San Saba River near Brady, Texas                    |
| 16     | 08145000 | Brady Creek at Brady, Texas                         |
| 17     | 08145100 | Brady Creek near Brady, Texas                       |
| 18     | 08145600 | San Saba River at San Saba, Texas                   |
| 19     | 08148300 | North Llano River near Junction, Texas              |
| 20     | 08150000 | Llano River near Junction, Texas                    |
| 21     | 08150200 | Llano River tribut. near London, Texas              |
| 22     | 08150700 | Llano River near Mason, Texas                       |
| 23     | 08150800 | Beaver Creek near Austin, Texas                     |
| 24     | 08150900 | Stone Creek tribut. near Art, Texas                 |
| 25     | 08151000 | Llano River near Castell, Texas                     |
| 26     | 08151300 | Johnson Creek near Valley Spring, Texas             |
| 27     | 08151500 | Llano River at Llano, Texas                         |
| 28     | 08152000 | Sandy Creek near Kingland, Texas                    |
| 29     | 08152700 | Little Flatrock Creek near Marble Falls, Texas      |
| 30     | 08128600 | Spring Creek near Fredericksburg, Texas             |
| 31     | 08129500 | Pedernales River near Fredericksburg, Texas         |
| 32     | 08130000 | Pedernales River at Stonewall, Texas                |
| 33     | 08131100 | Cave River at Stonewall, Texas                      |
| 34     | 08153500 | Pedernales River near Johnson City, Texas           |
| 35     | 08154000 | Pedernales River near Spicewood, Texas              |
| 36     | 08154700 | Barton Creek at Loop 360 near Austin, Texas         |
| 37     | 08155200 | Blanco Creek at SH 171 near Oak Hill, Texas         |
| 38     | 08155300 | Blanco Creek at SH 171 near Oak Hill, Texas         |
| 39     | 08158700 | Onion Creek near Driftwood, Texas                   |
| 40     | 08158810 | Blanco below FM Road 1826 near Driftwood, TX        |
| 41     | 08158840 | Slaughter Creek at FM Road 1826 near Austin, Texas  |
| 42     | 08159800 | Fry Branch near Oak Hill, Texas                     |
| 43     | 08159150 | Widganger Creek near Pflugerville, Texas            |
| 44     | 08160300 | North Fork Guadalupe River near Hunt, Texas         |
| 45     | 08160500 | Guadalupe River at Hunt, Texas                      |
| 46     | 08160600 | Johnson Creek near Ingram, Texas                    |
| 47     | 08160200 | Guadalupe River at Kerrville, Texas                 |
| 48     | 08160300 | Turkey Creek near Kerrville, Texas                  |
| 49     | 08167000 | Guadalupe River at Comfort, Texas                   |
| 50     | 08167500 | Guadalupe River near Spring Branch, Texas           |
| 51     | 08167600 | Rebecca Creek near Spring Branch, Texas             |
| 52     | 08171000 | Blanco River at Winberryville, Texas                |
| 53     | 08173100 | Blanco River near Kyle, Texas                       |
| 54     | 08180900 | Cibola Creek near Big Bend, Texas                   |
| 55     | 08184000 | Cibola Creek near Bulverde, Texas                   |